

Volcano Deformation and Physics Based Eruption Models

Paul Segall

Stanford University

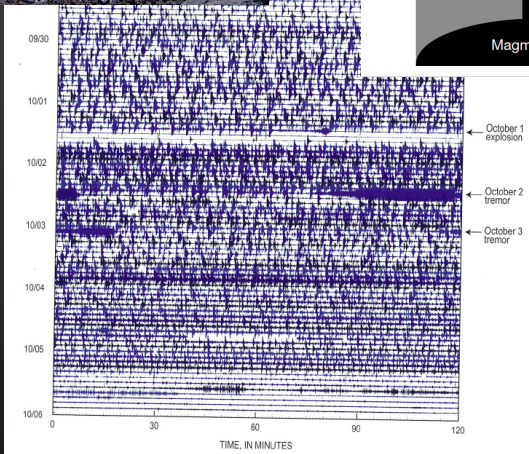
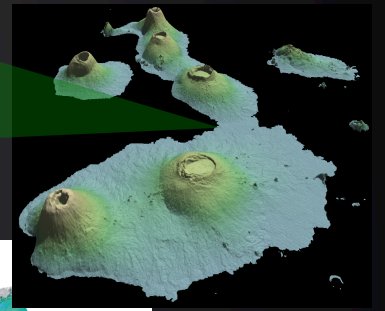
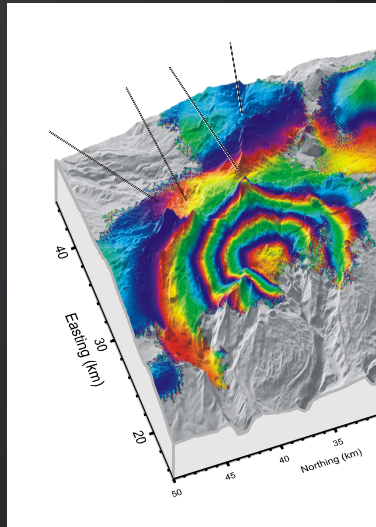
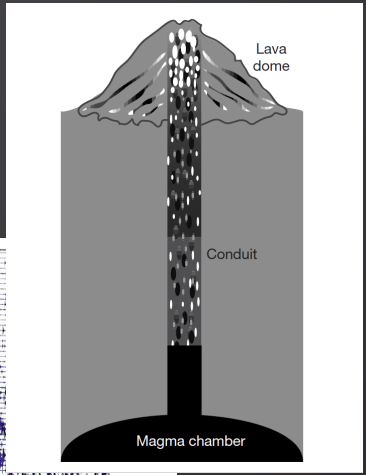
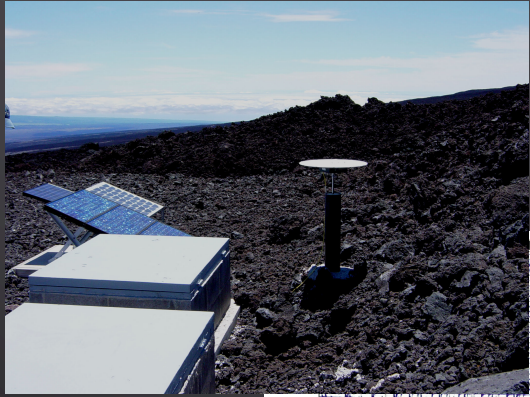
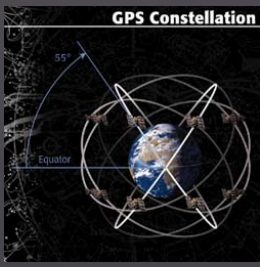
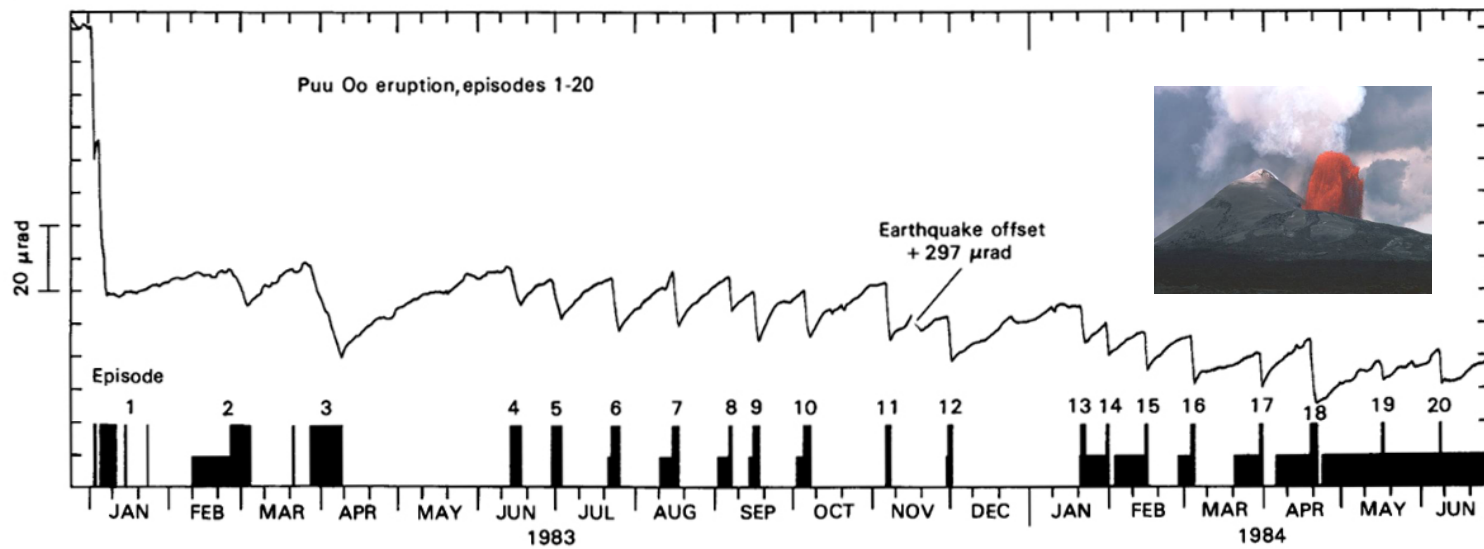
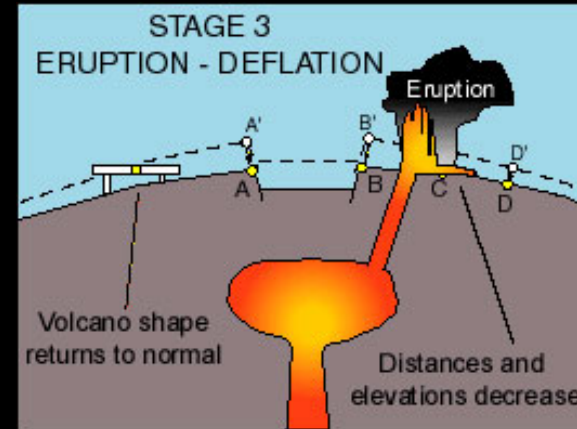
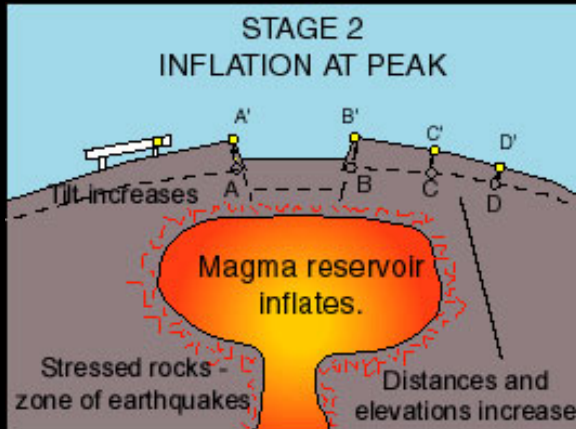


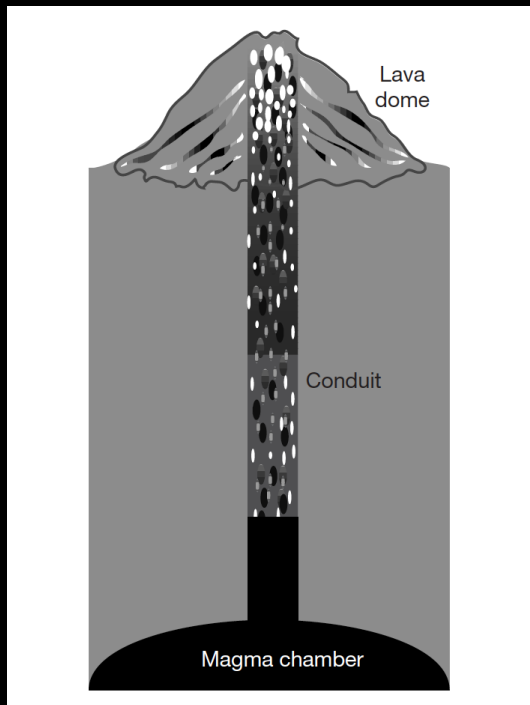
Figure 8. Seismic record from station JUN (6.5 km southeast of vent; fig. 1) from September 27 through October 5, 2004, showing seismicity trends during the vent-clearing phase.



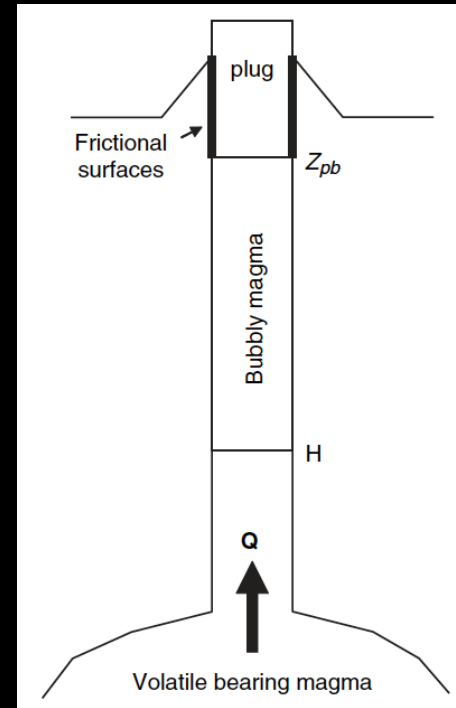
Inflation Deflation Cycles



Physics Based Eruption Models



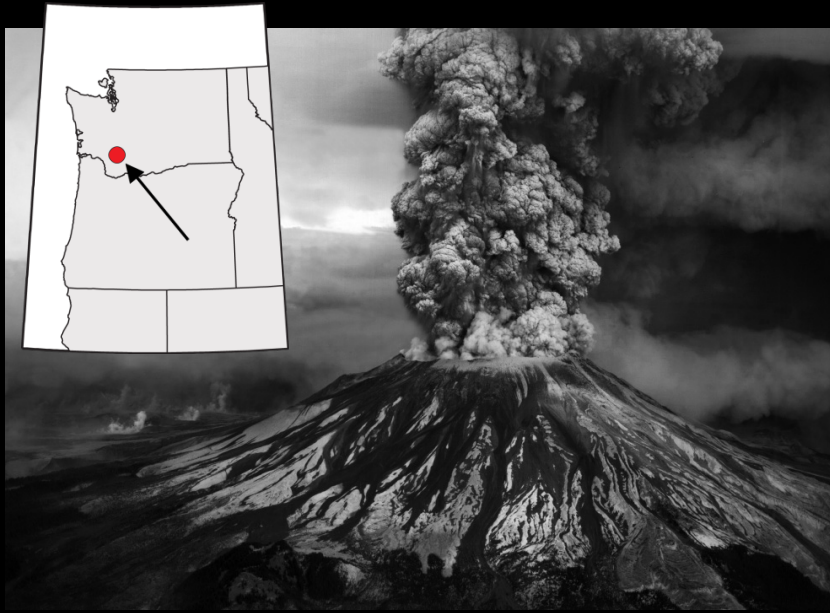
Melnik and Sparks, 1999, Nature



Lensky et al 2009, Geol Soc Lond

[Stasiuk et al., 1993; Ramos, 1995; Jaupart, 1996; Melnik and Sparks, 1999; 2002, 2005; 2006, Mastin and Ghiorso, 2000; Maeda, 2000; Massol et al., 2001; Huppert and Woods, 2002; Woods and Huppert, 2003; Barmin et al., 2002; Proussevitch and Sahagian, 2005; Starostin et al., 2005; Collier and Neuberg, 2006; Mason et al., 2006; Costa et al., 2007a, b; de' Michieli Vitturia et al., 2008; Lensky et al. 2008; Hautmann et al., 2009; Mastin et al., 2008; 2009]

Mount St. Helens, Washington

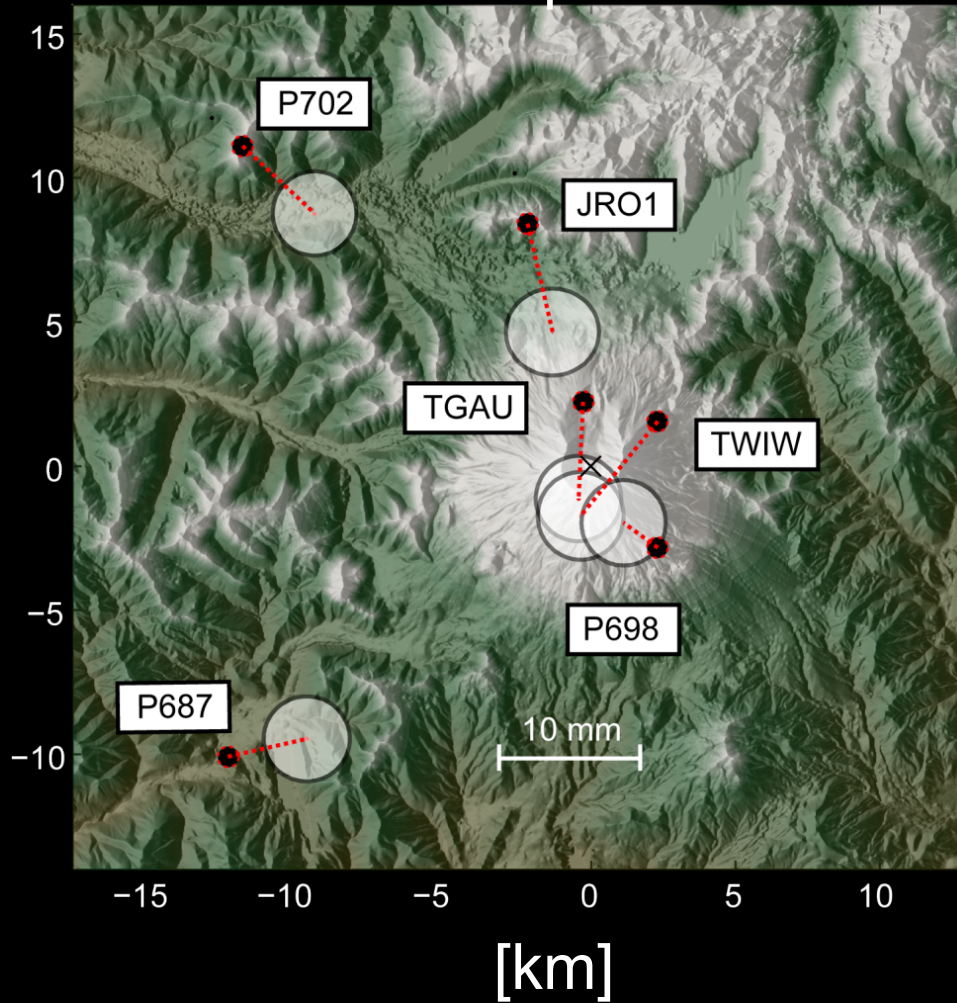


- Catastrophic eruption 1980
- Effusive eruption 2004-2008

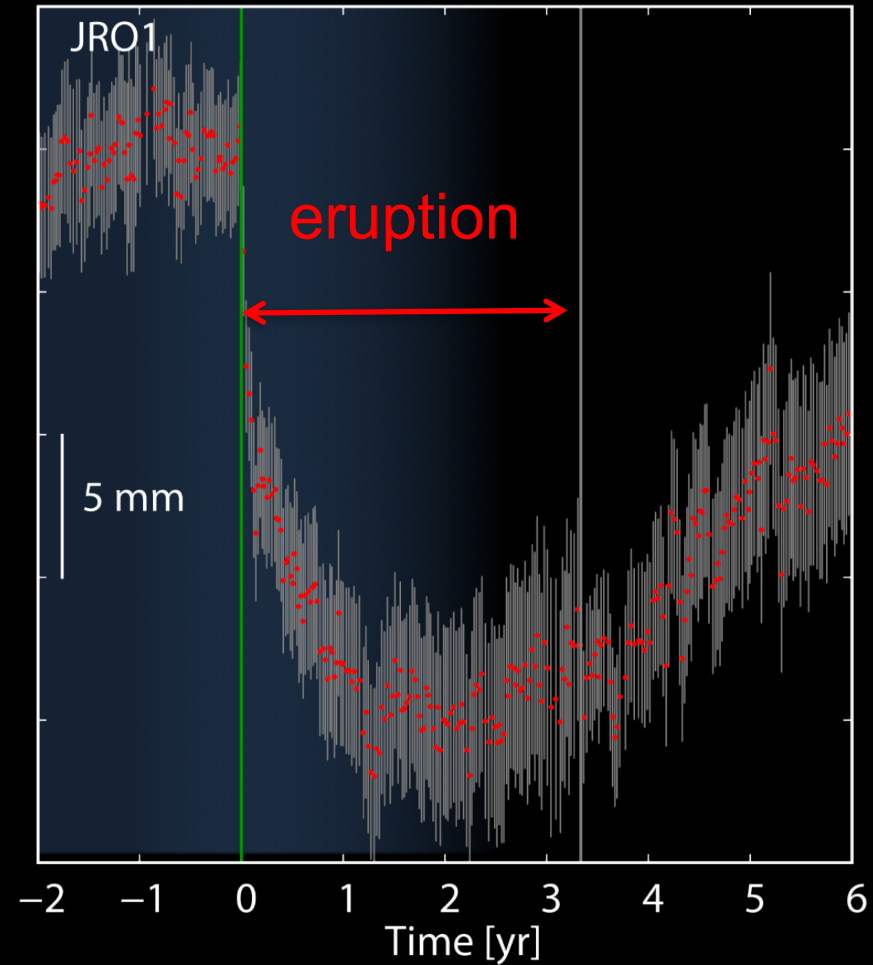


Mount St Helens Dome Forming Eruption 2004-2008

Net Displacements

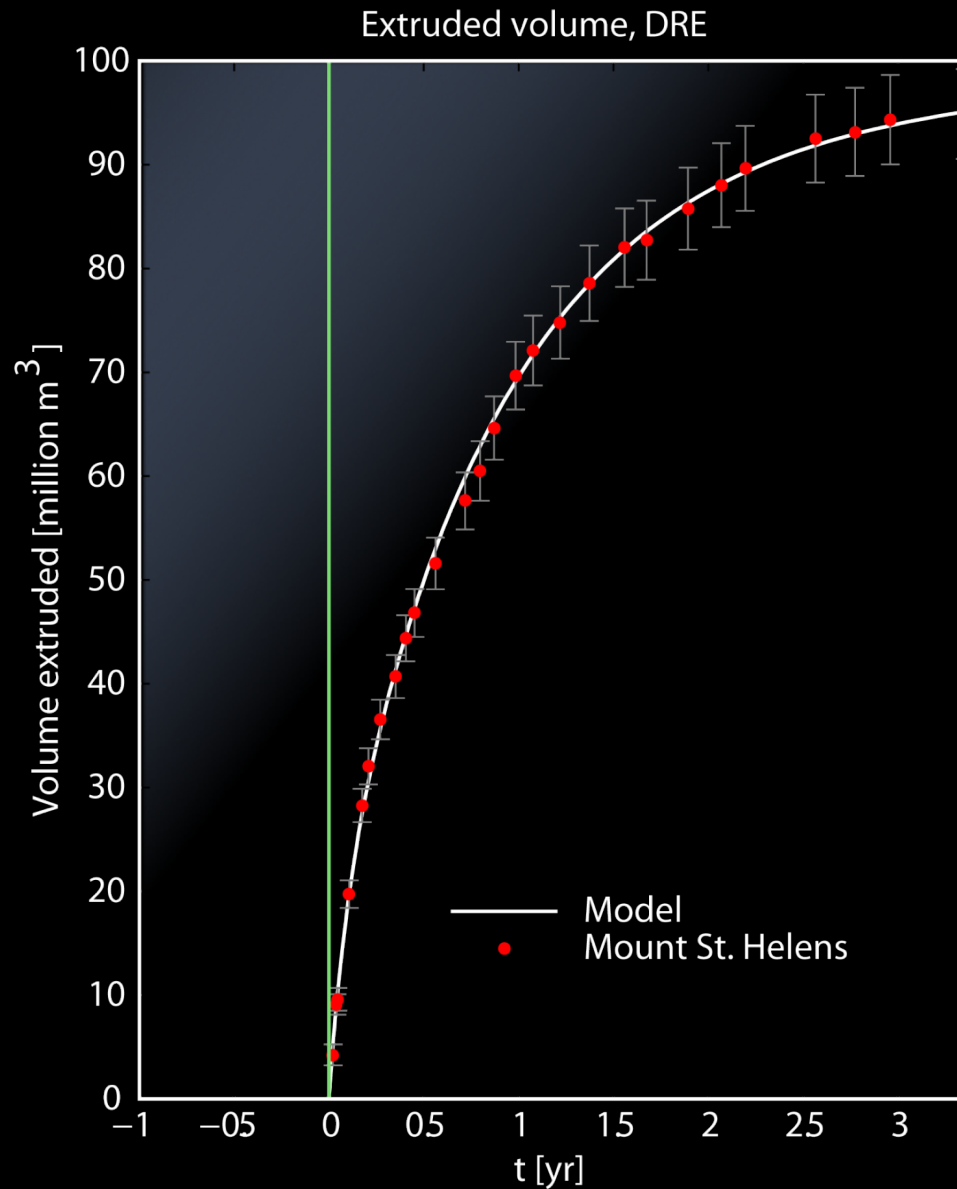
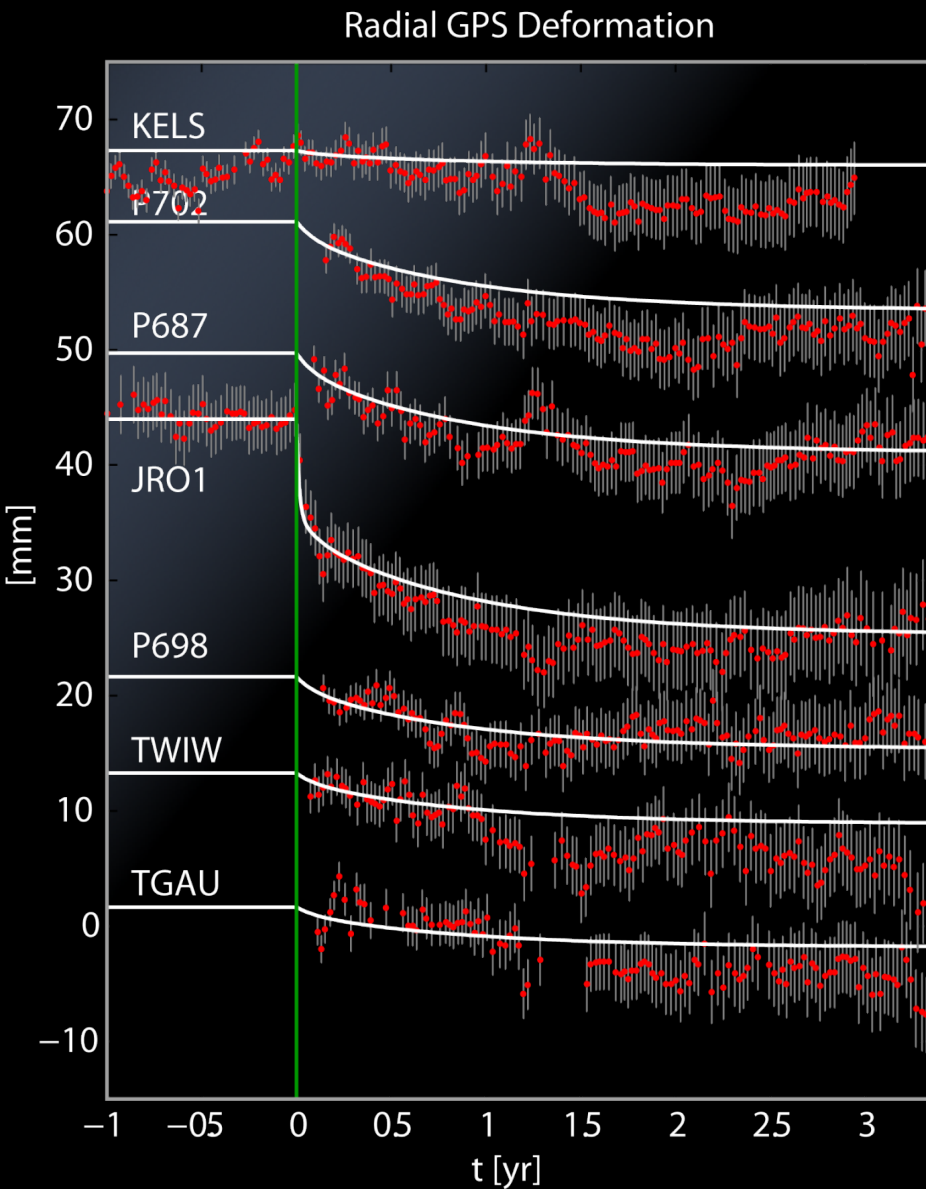


JRO1 Radial Time Series

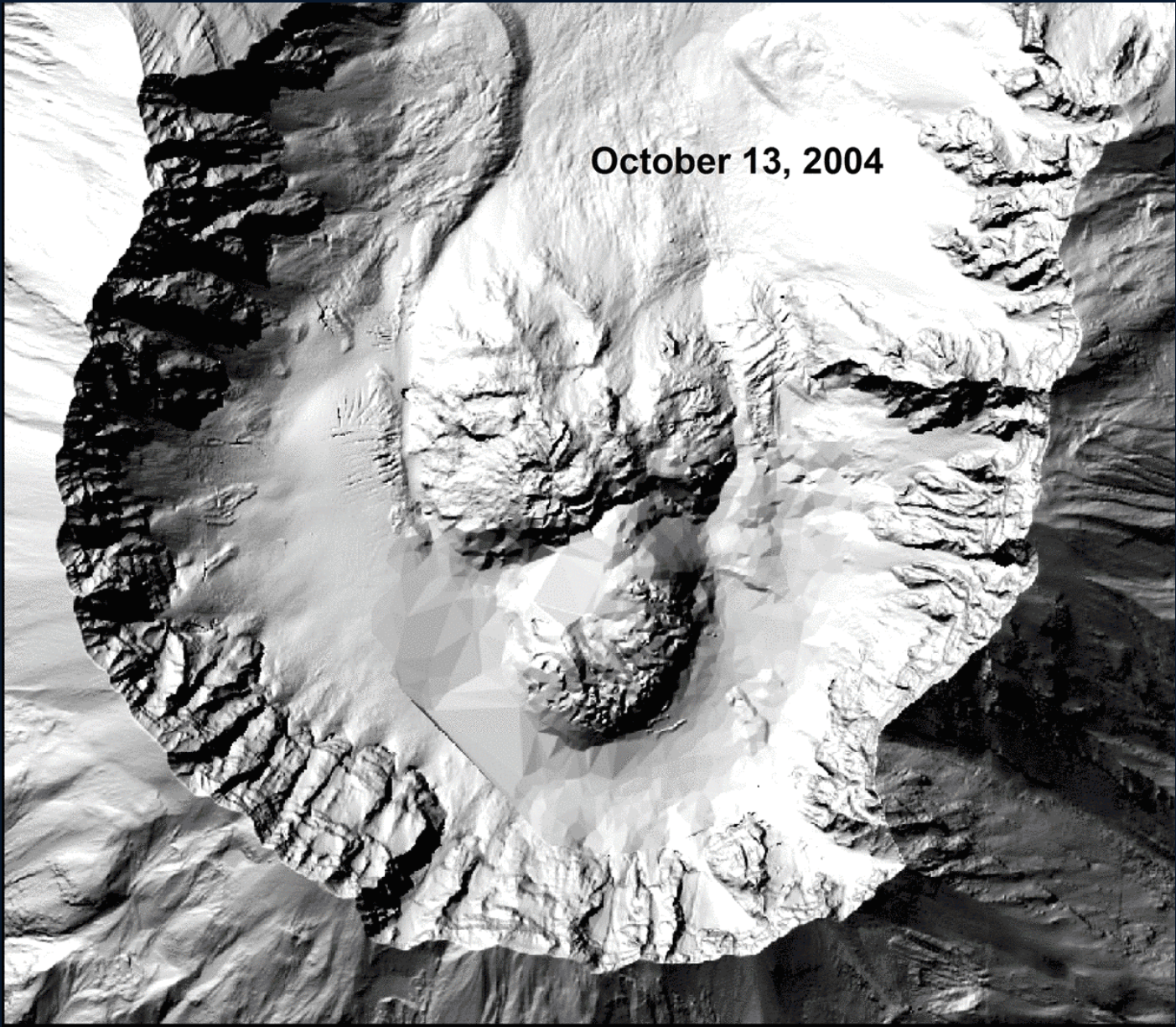


Lisowski et al. [2008]

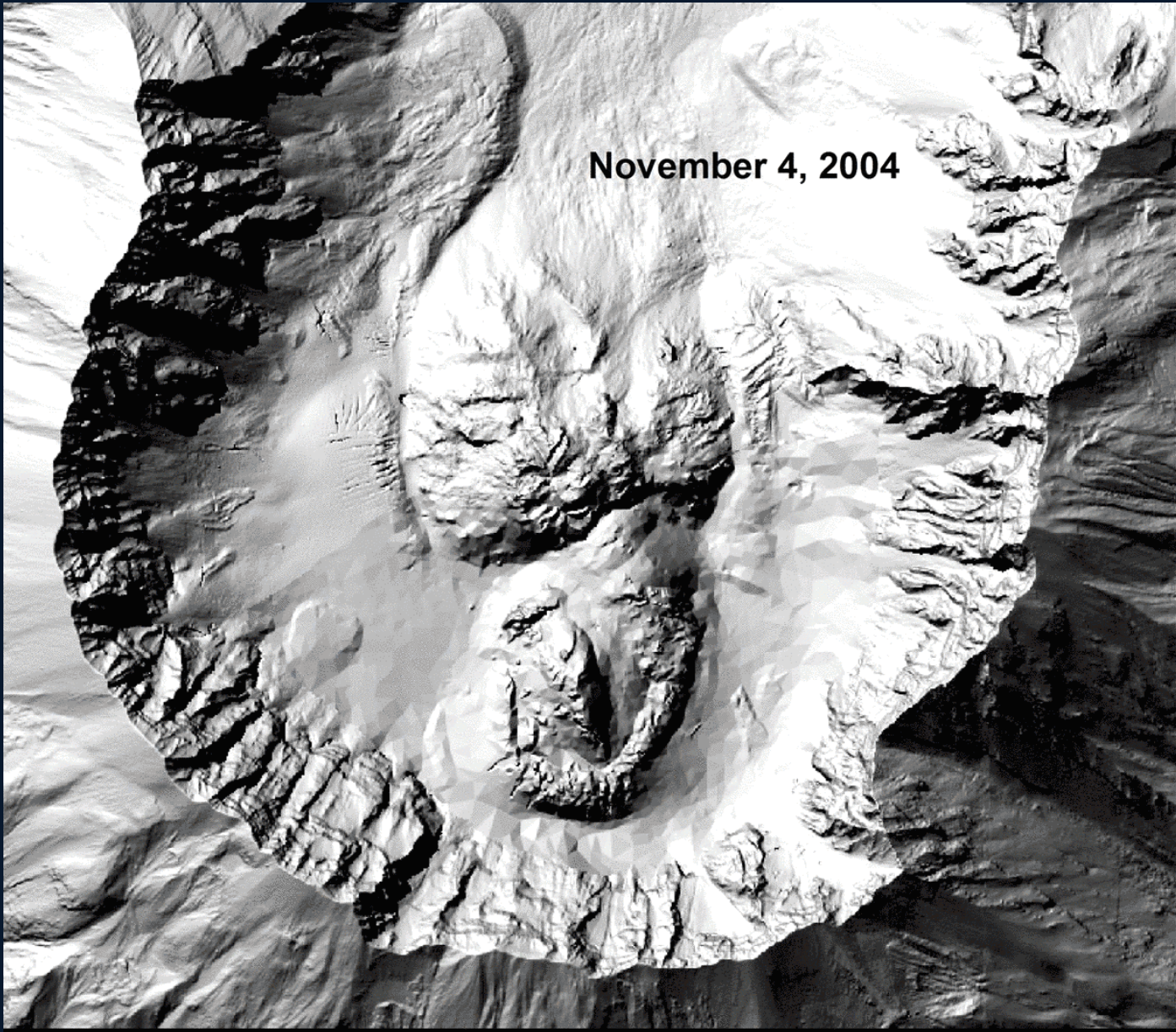
Model Fits both GPS and Extrusion Data



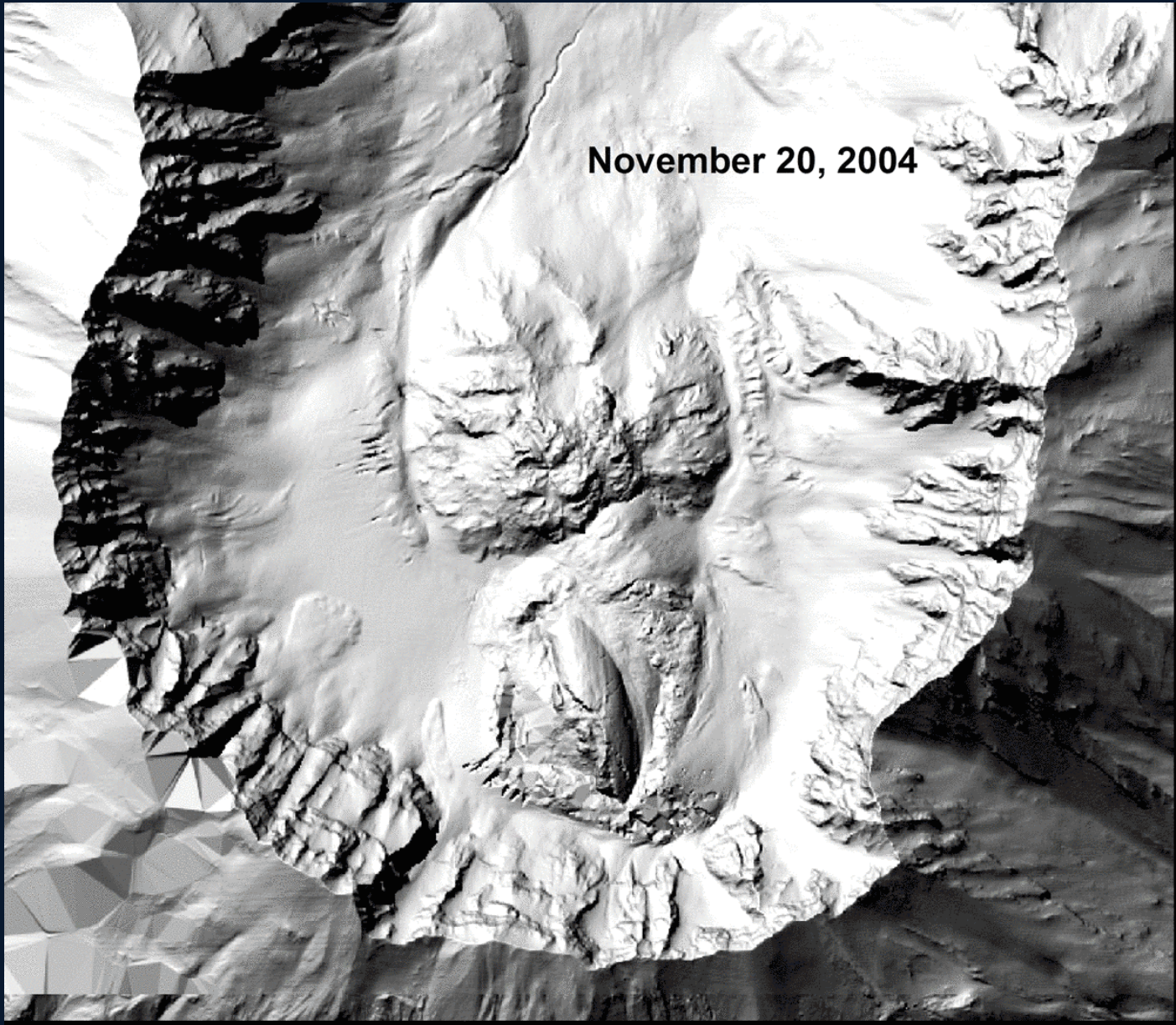
October 13, 2004



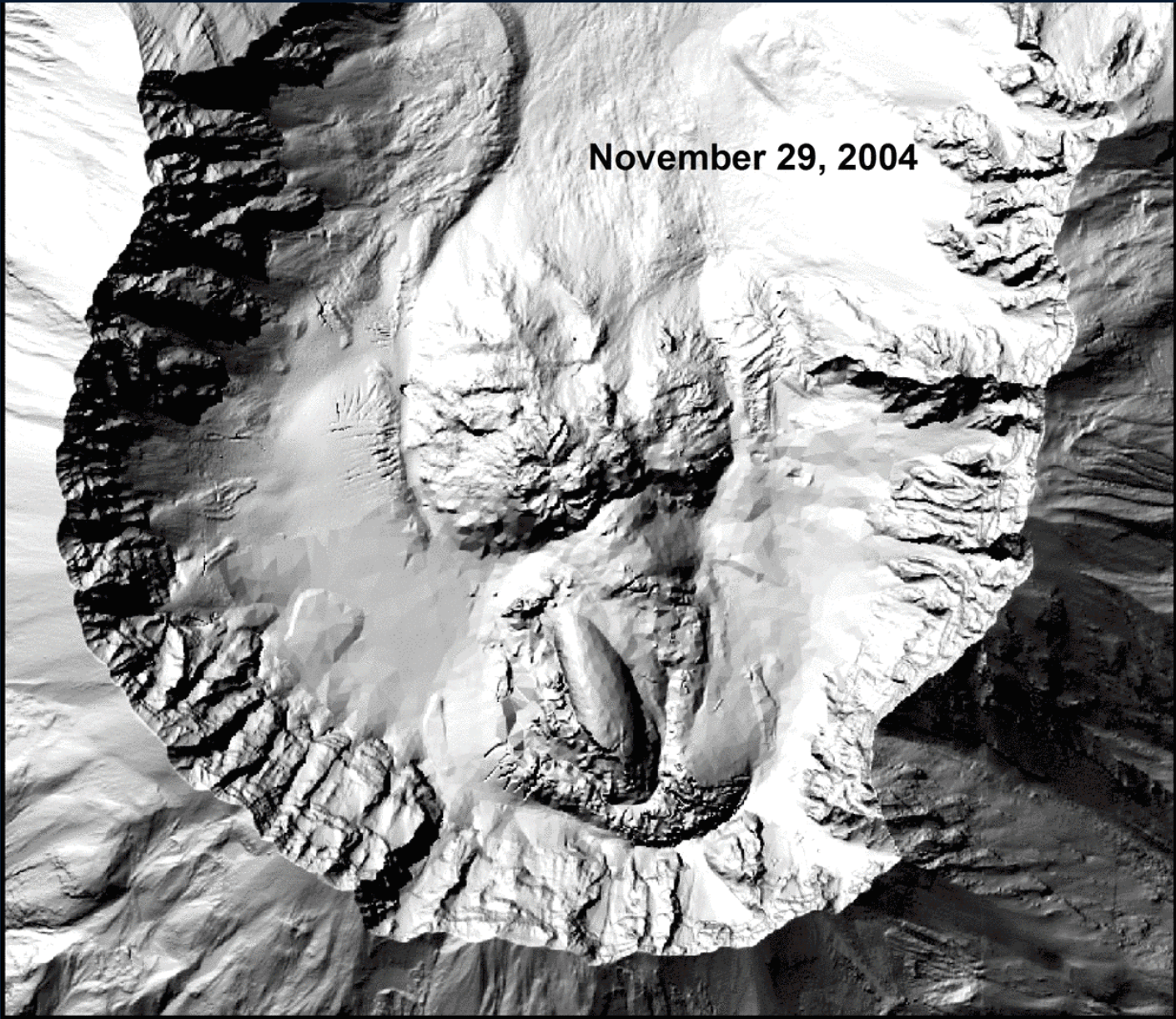
November 4, 2004



November 20, 2004



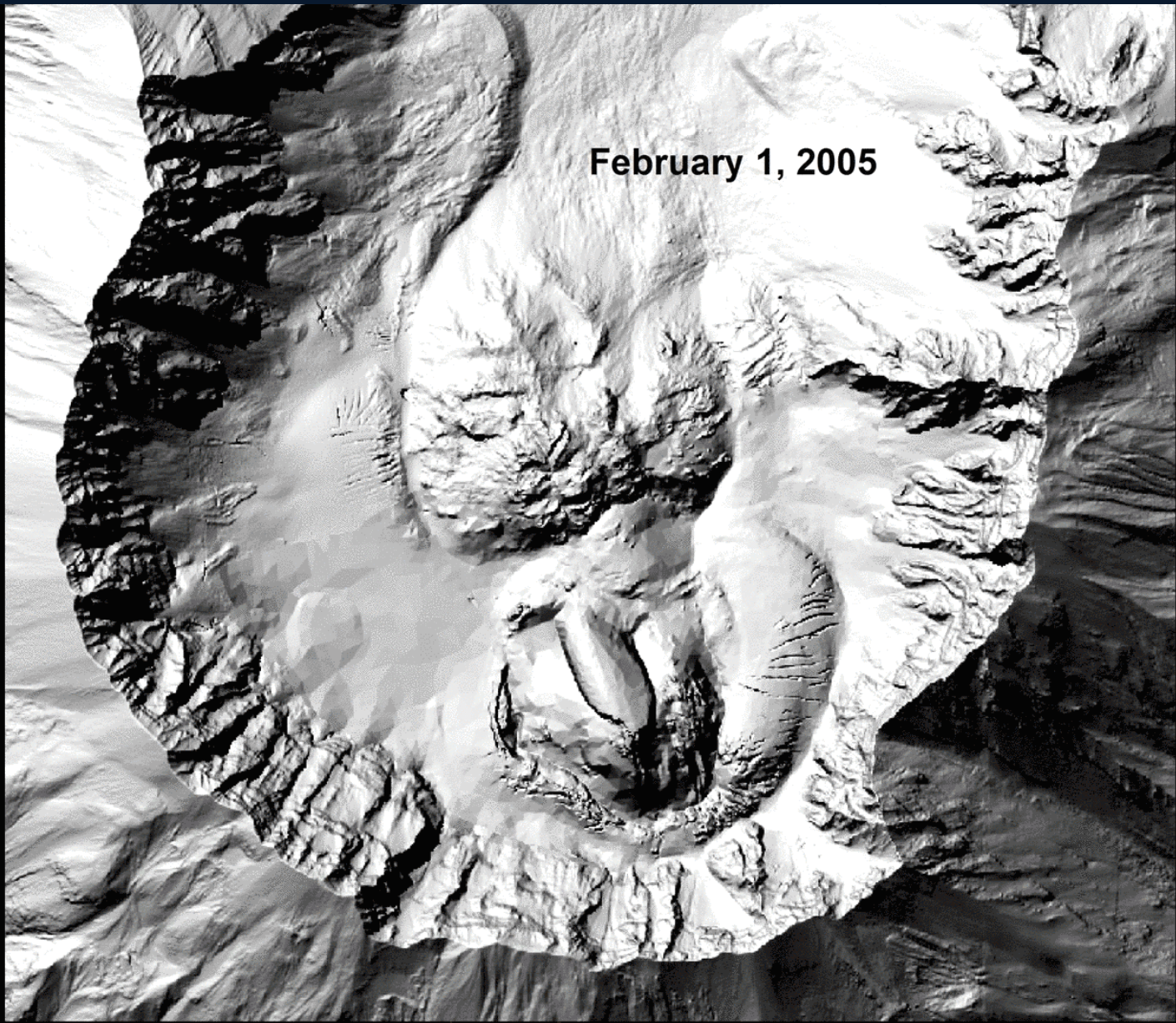
November 29, 2004



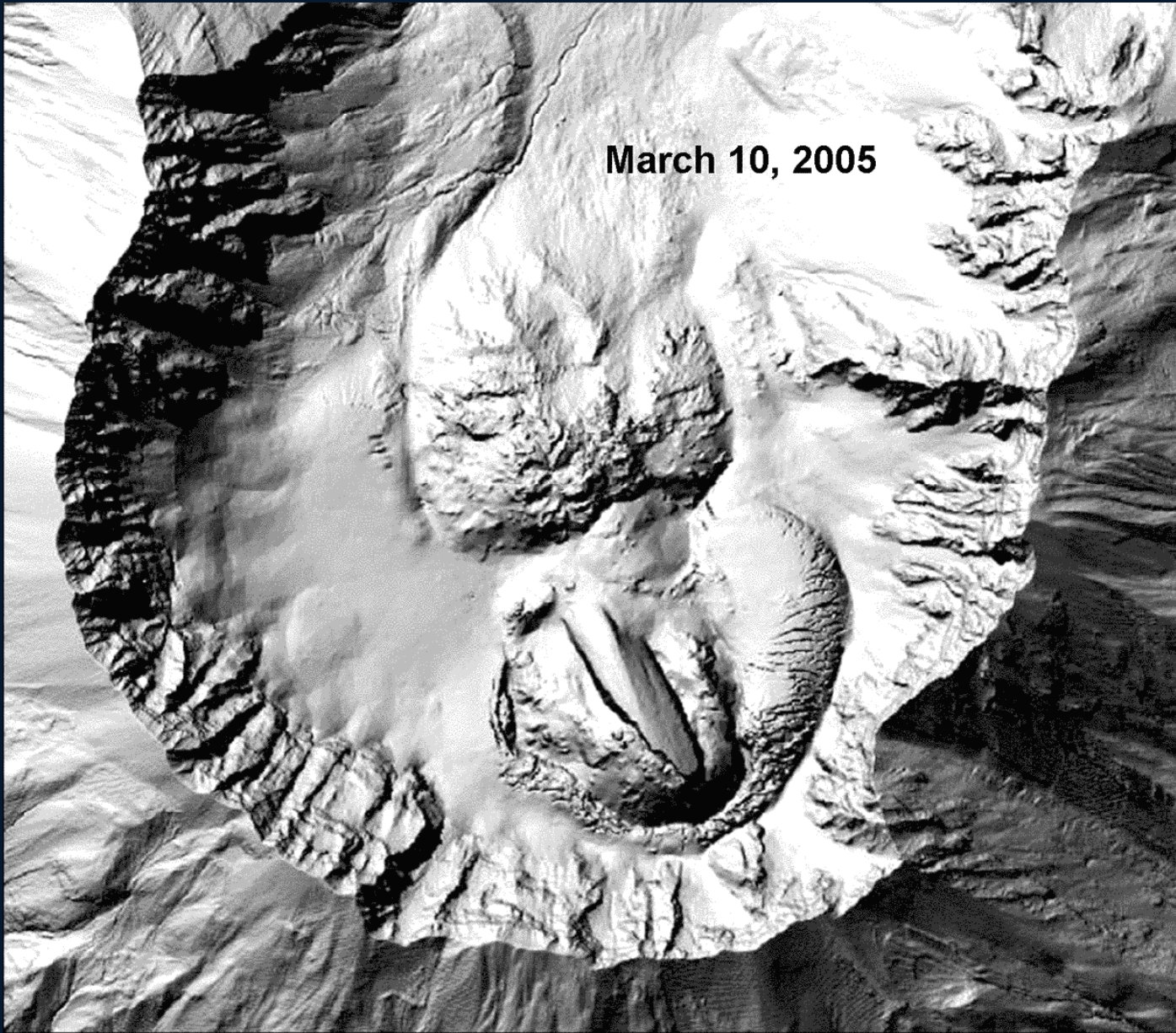
December 11, 2004



February 1, 2005



March 10, 2005



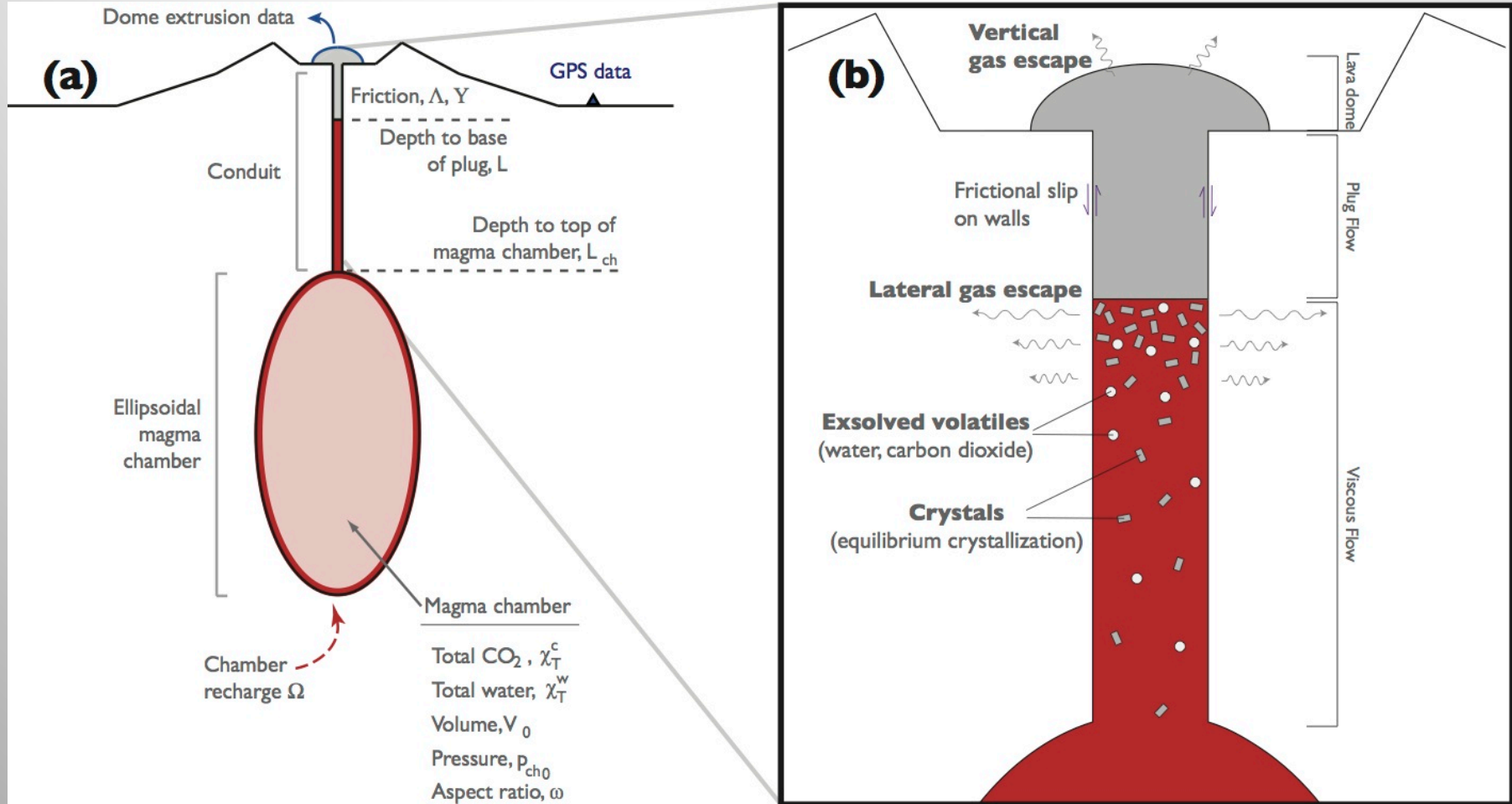
Physics-based Volcano Deformation



Kyle Anderson



YingQi Wong



Governing equations

Crystallization depth (fixed)

Chamber Equations

Compressibility

$$\beta_m = \frac{1}{\rho_{cc}} \frac{\partial \rho_{cc}}{\partial p_{cc}}$$

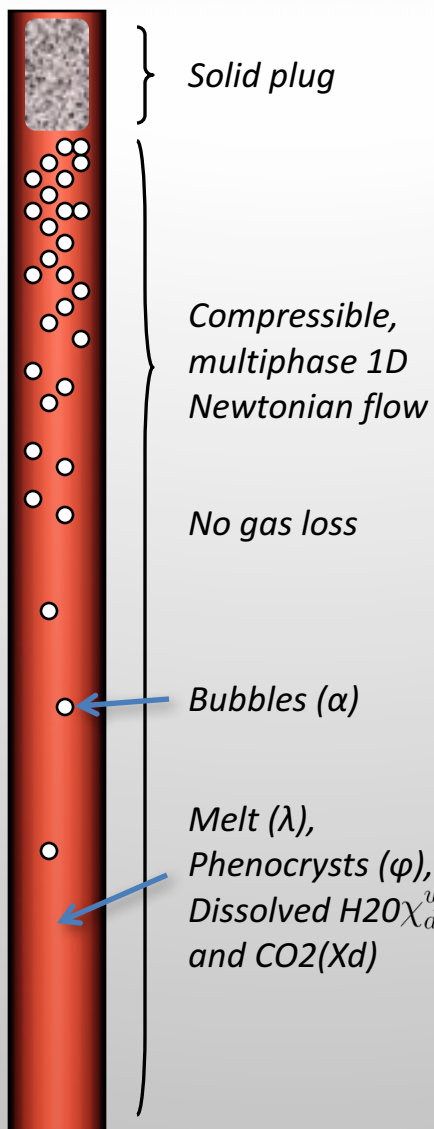
$$\beta_{ch} = \frac{1}{V} \frac{\partial V}{\partial p_{cc}}$$

Evolution of pressure

$$\frac{dp_{ch}}{dt} = \frac{q_{in} - q_{out}}{V_0 (\beta_m + \beta_{ch})}$$

Chamber influx

$$q_{in} = \Omega (p_{deep} - p_{ch})$$



Conduit Equations

Momentum & mass conservation

$$0 = \frac{\partial p}{\partial z} + \rho g + \frac{8\eta v}{R^2}$$

$$\frac{\partial \rho}{\partial t} = -\frac{\partial}{\partial z} (\rho v)$$

Magma viscosity

$$\eta = \eta_\chi (\chi_d^w, T) \eta_\phi (\phi^v)$$

Plug friction

$$\tau_p = a \bar{\sigma} \operatorname{arcsinh} \left(\frac{v}{2v_r} \exp \frac{f_0}{a} \right)$$

Density

$$\rho^* = \left[\frac{1 - \phi}{1 + \chi_e + \chi_d} \left(\frac{\chi_e}{\rho_\alpha} + \frac{\chi_d}{\rho_c} + \frac{1}{\rho_\lambda} \right) + \frac{\phi}{\rho_\phi} \right]^{-1}$$

Solubility of H₂O and CO₂ in melt

Mole fraction H₂O and CO₂ in gas phase

Pressure and temperature

$$\text{H}_2\text{O: } \chi_d^w = \chi_T^w - \chi_e^w = S^w(m^w, m^c, p, T)$$

$$\text{CO}_2: \chi_d^c = \chi_T^c - \chi_e^c = S^c(m^w, m^c, p, T)$$

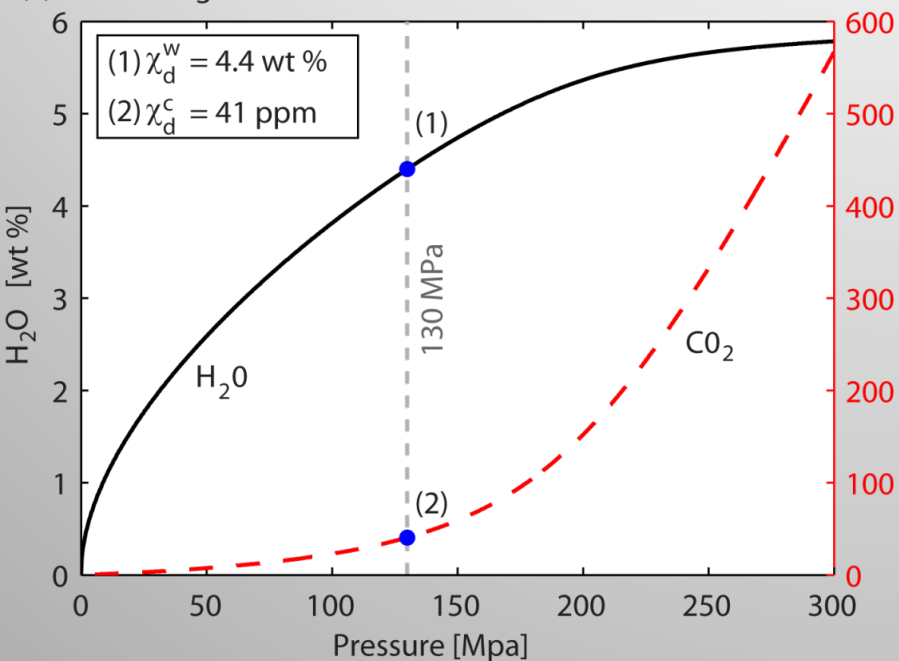
Dissolved volatiles

Total volatiles

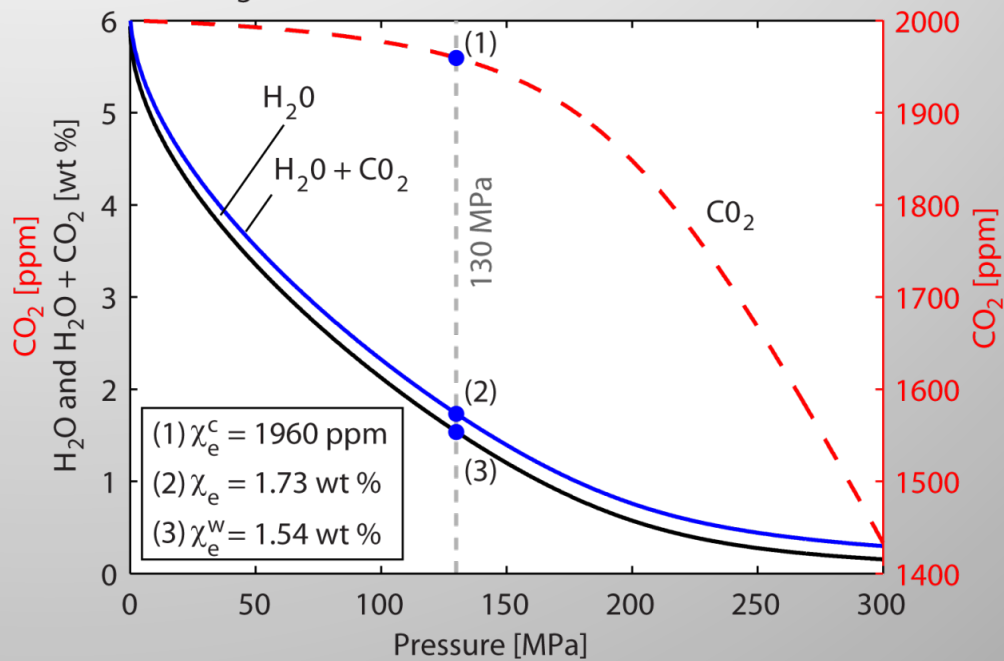
Exsolved volatiles

Empirical solubility relationship from *Liu et al.* [2005]

(a) Dissolved gases



(b) Exsolved gases



Bayes' Theorem

$$P(m|d) \propto P(d|m) P(m)$$

Posterior
Probability

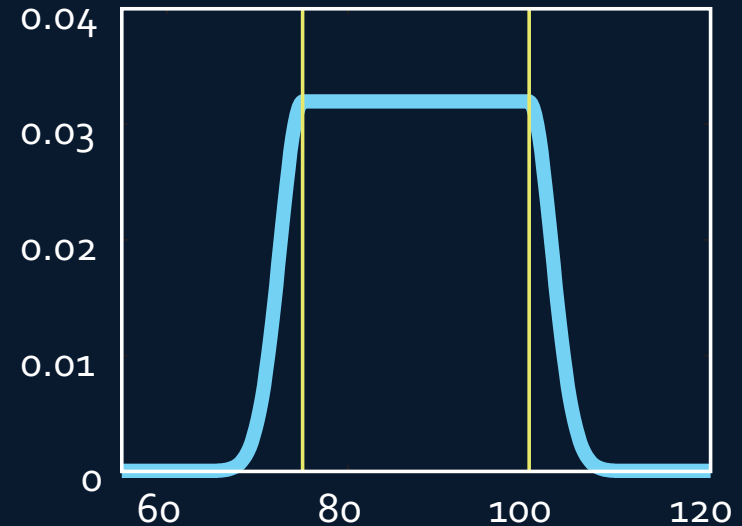
Data
Likelihood

Prior
Probability

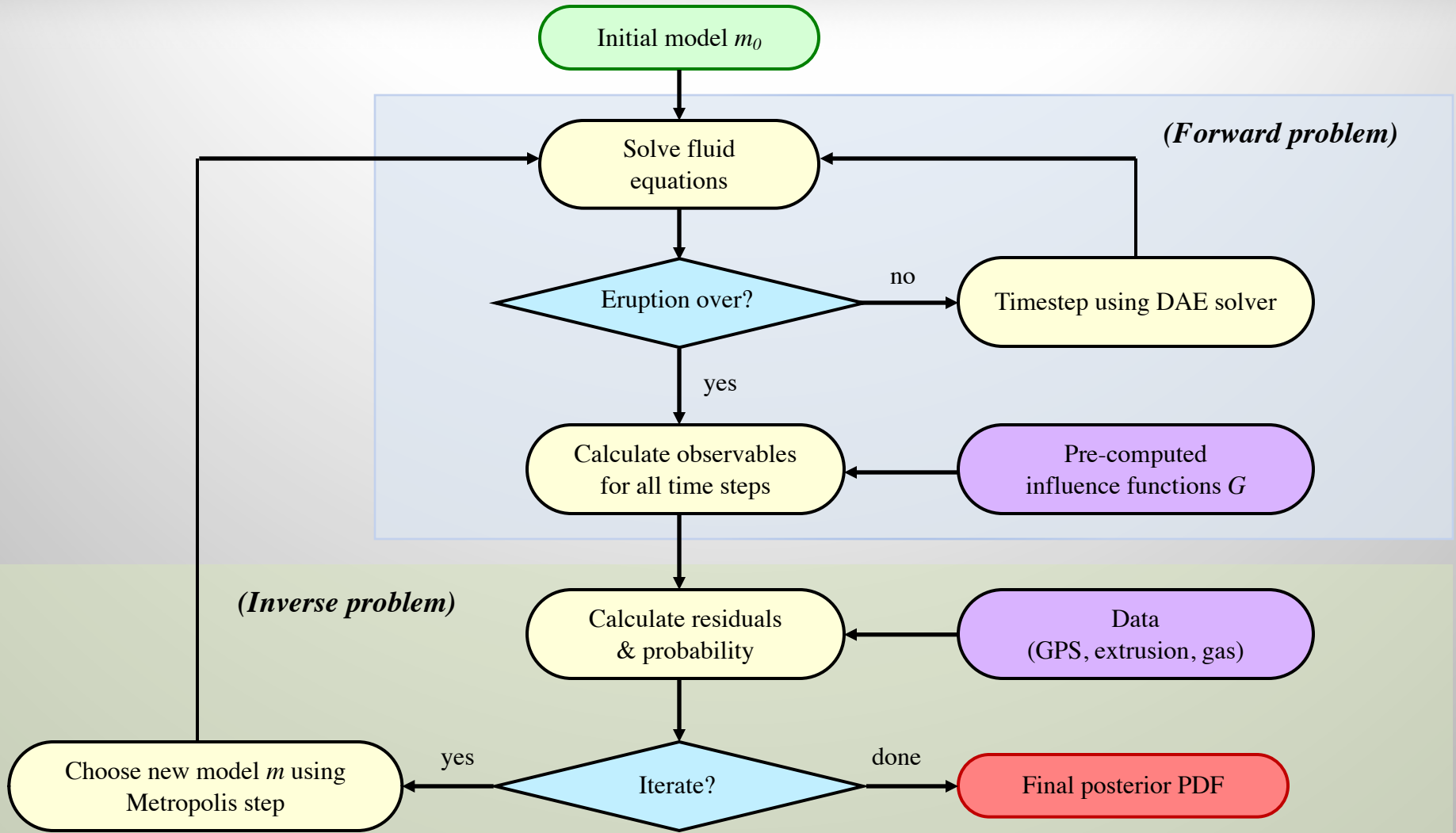
m: model
d: data

Use Markov Chain Monte Carlo (MCMC) method to characterize posterior probability distributions

Chamber Pressure
MPa



Incorporation into Markov-Chain Monte Carlo inversion

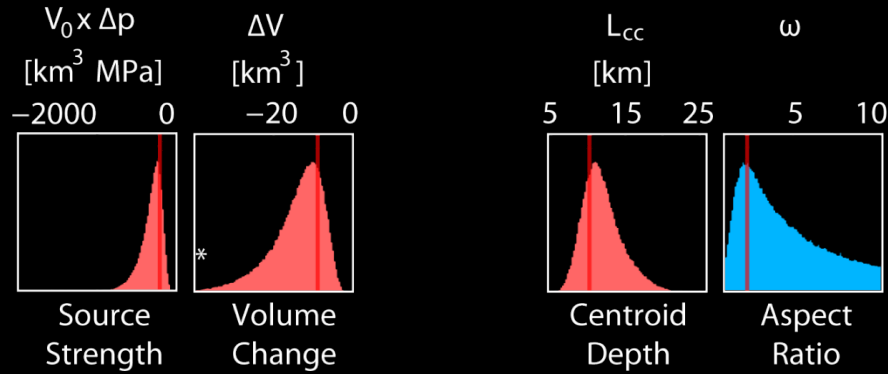


$$P(\mathbf{d}_1, \mathbf{d}_2, \dots, \mathbf{d}_K | \mathbf{m}) =$$

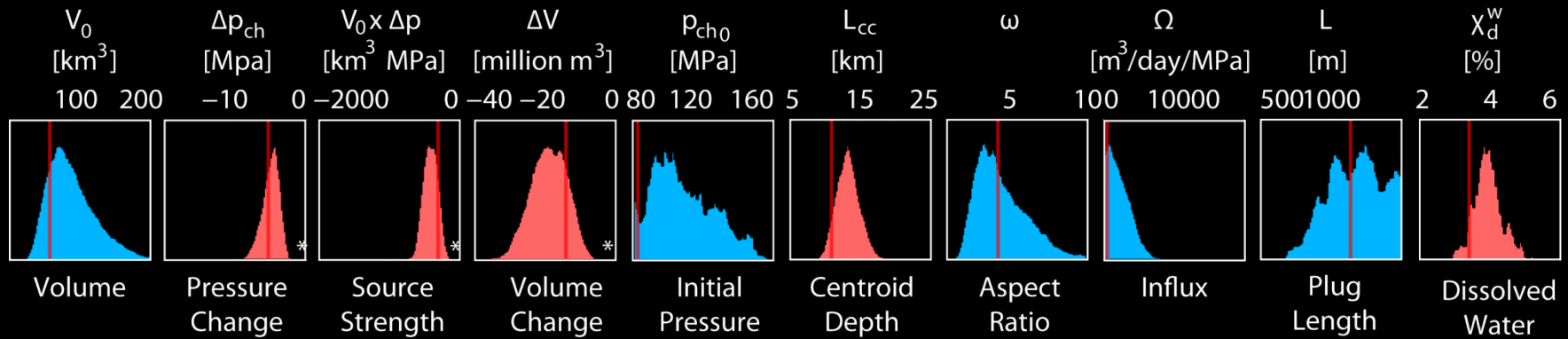
$$\prod_{k=1}^K \left\{ (2\pi\gamma_k^2)^{-N_k/2} |\Sigma_k|^{-1/2} \exp\left(-\frac{1}{2\gamma_k^2} \mathbf{r}_k^T \Sigma_k^{-1} \mathbf{r}_k\right) \right\}$$

Kinematic vs. Physics-Based Inversion

Kinematic

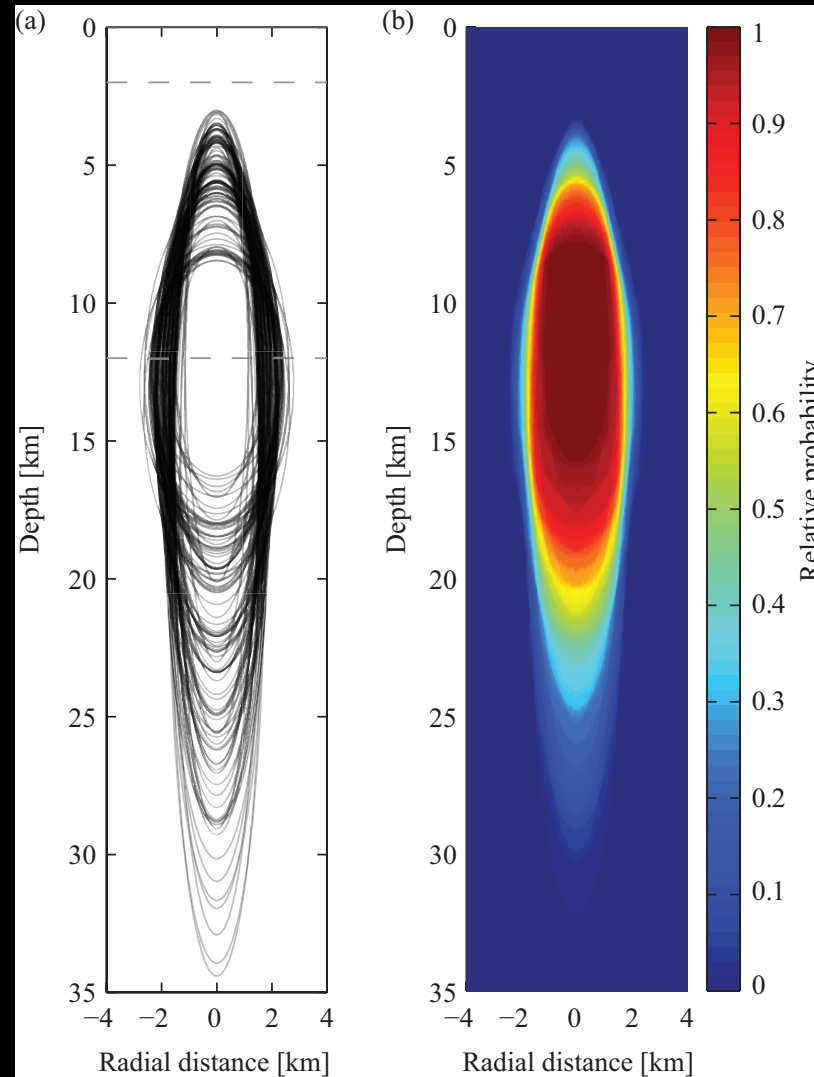


Physics-based

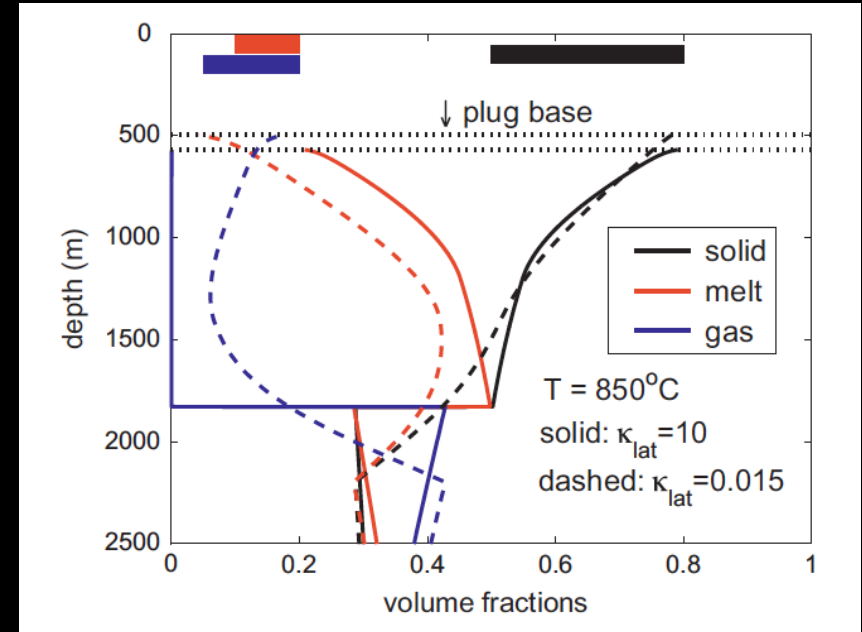
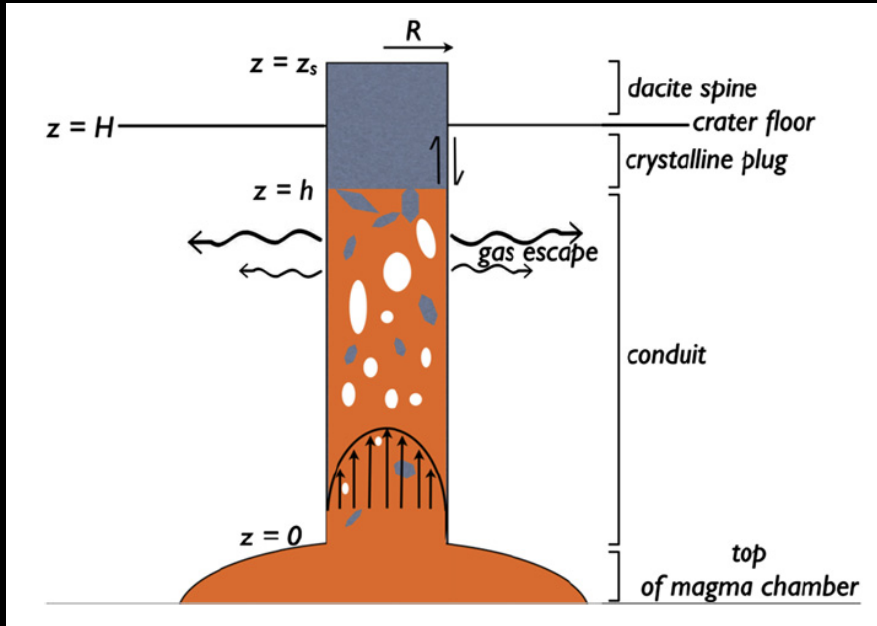


Anderson and Segall, JGR 2013

Summary of Magma Chamber Geometry

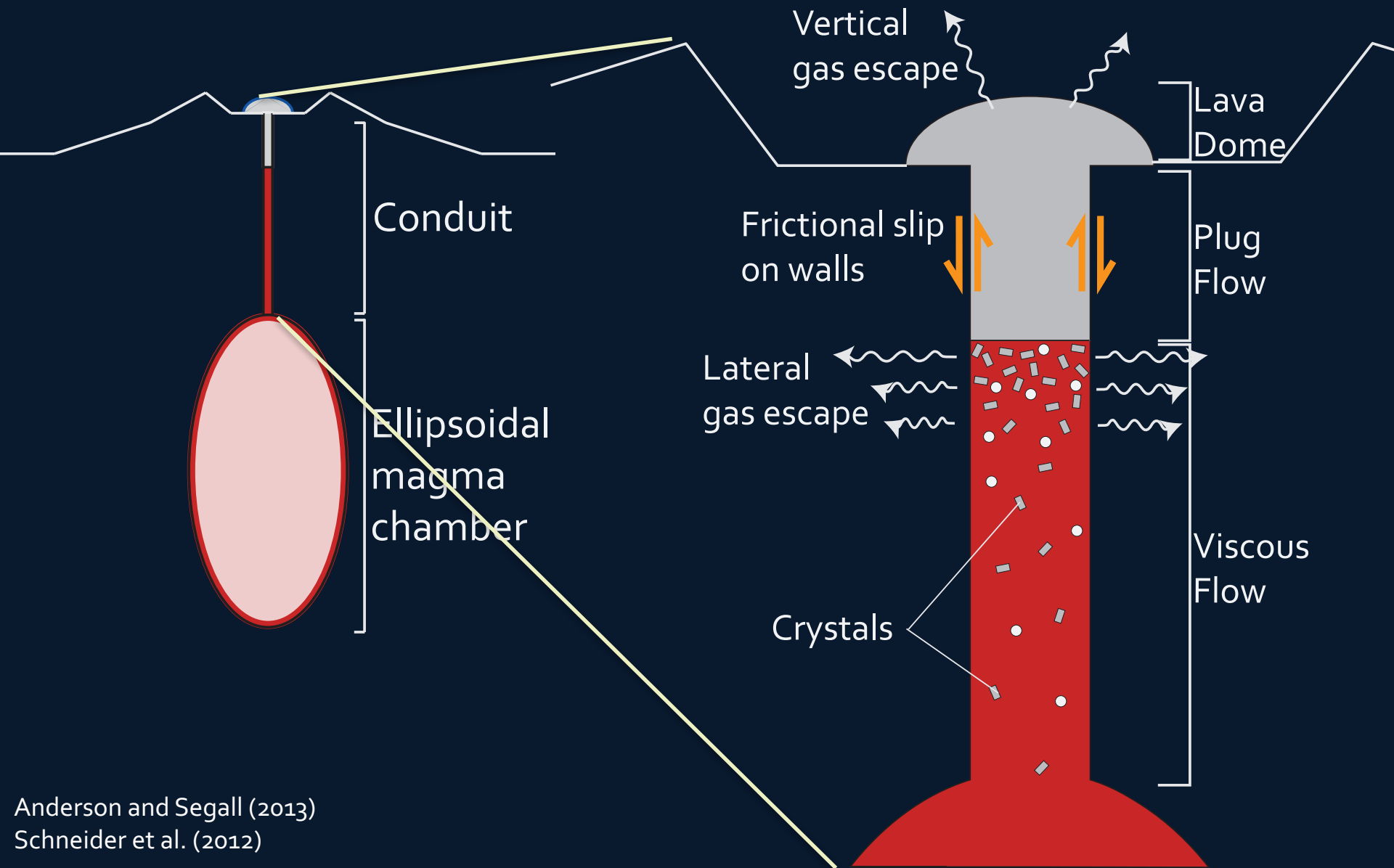


Schneider, Rempel, Cashman (2012)



- Steady state eruption model
- Equilibrium crystallization based on MELTS
- “Membrane diffusion” gas loss

Physics-based model: Geometry and Setup

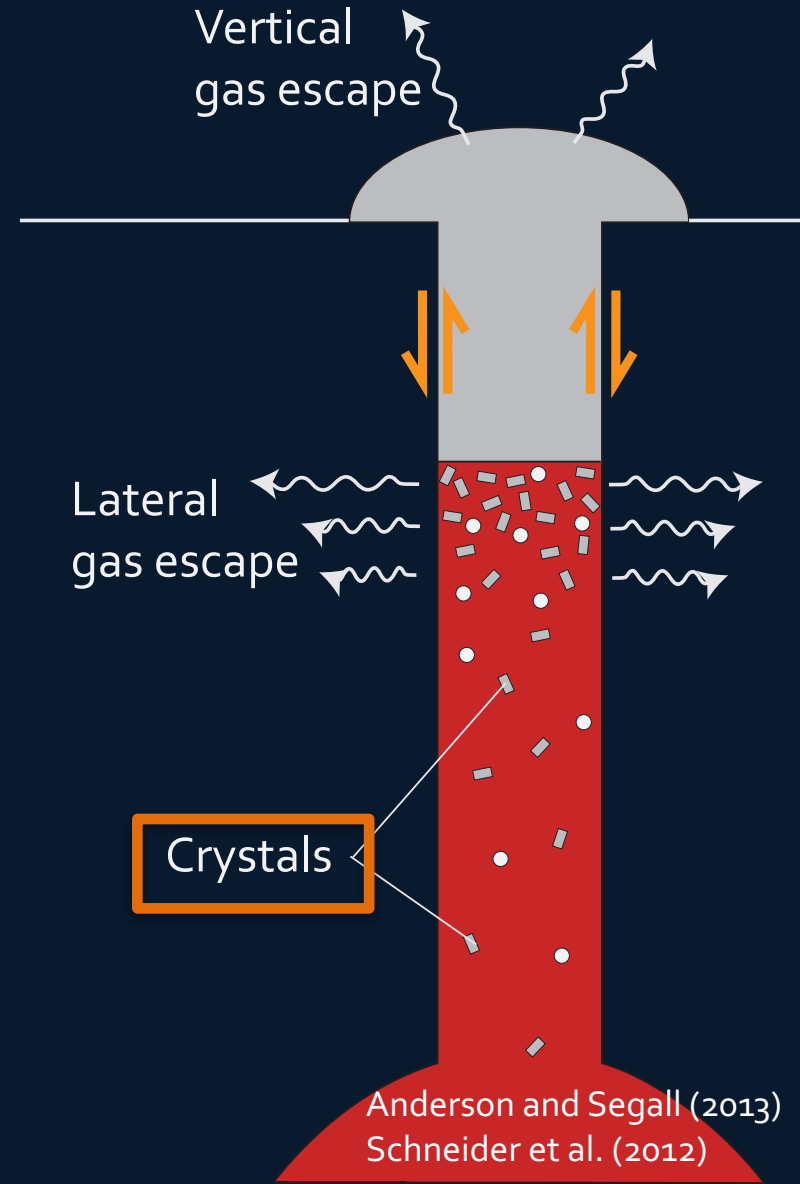
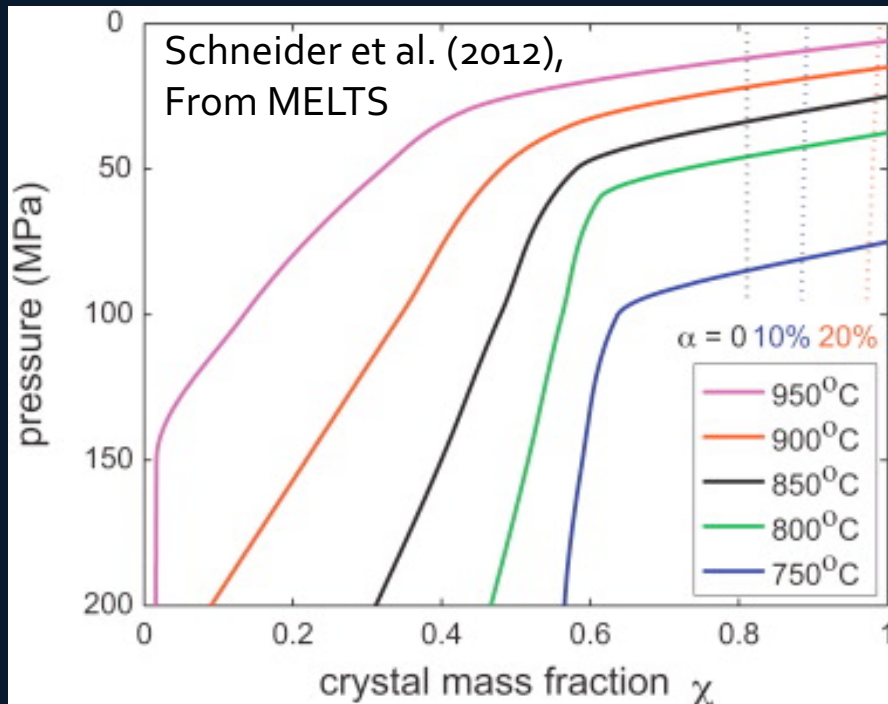


Equilibrium crystallization

Depressurization



Crystallization
(Assumed isothermal)



Equilibrium crystallization

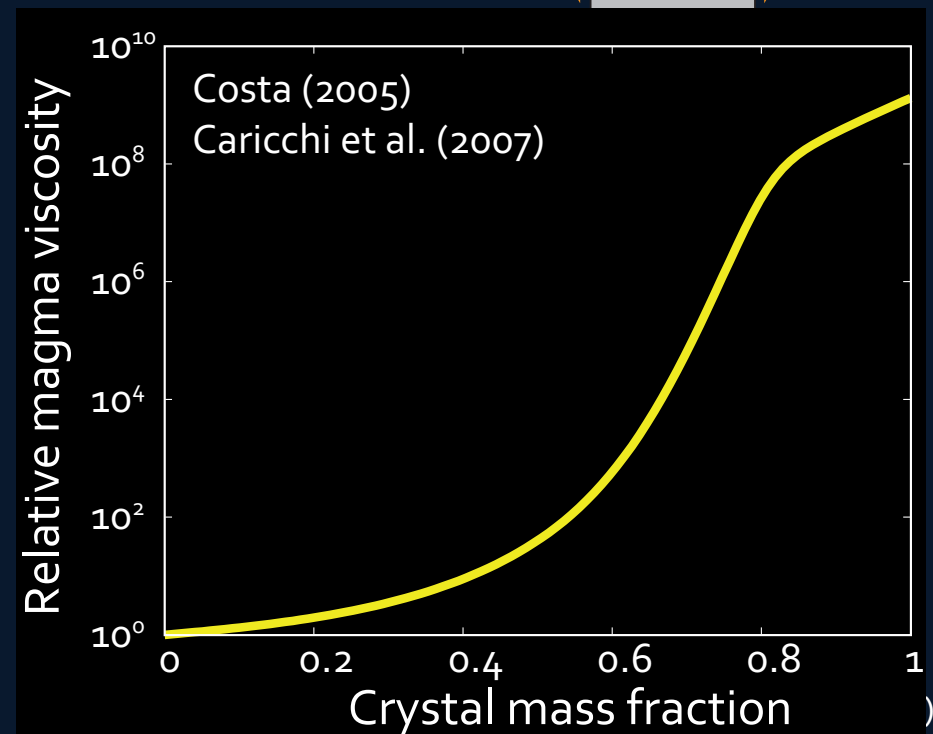
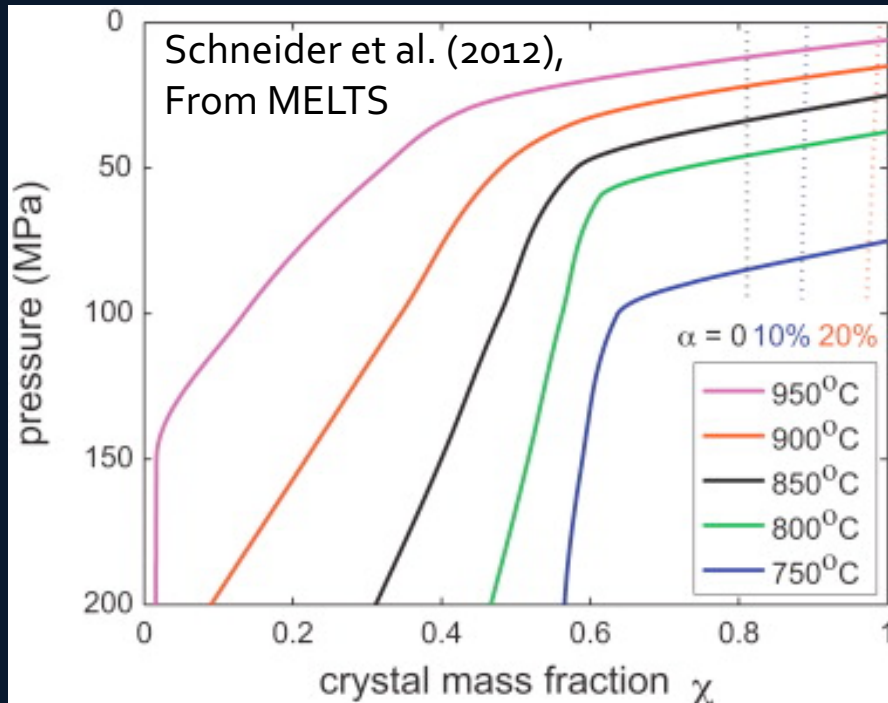
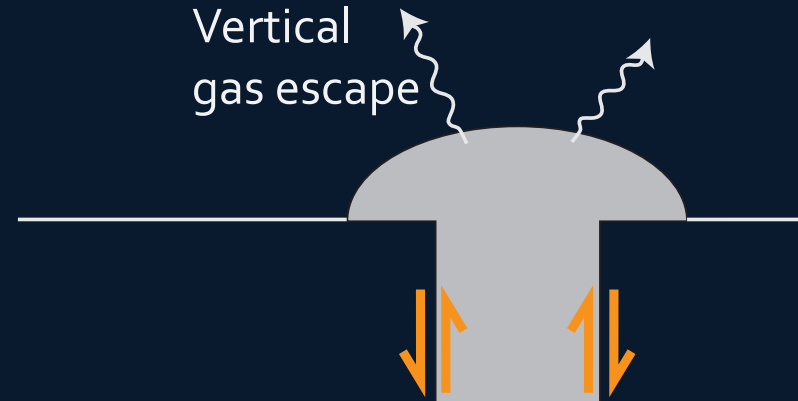
Depressurization



Crystallization



Increase in viscosity



Schneider et al. (2012)

Viscous flow to plug flow transition

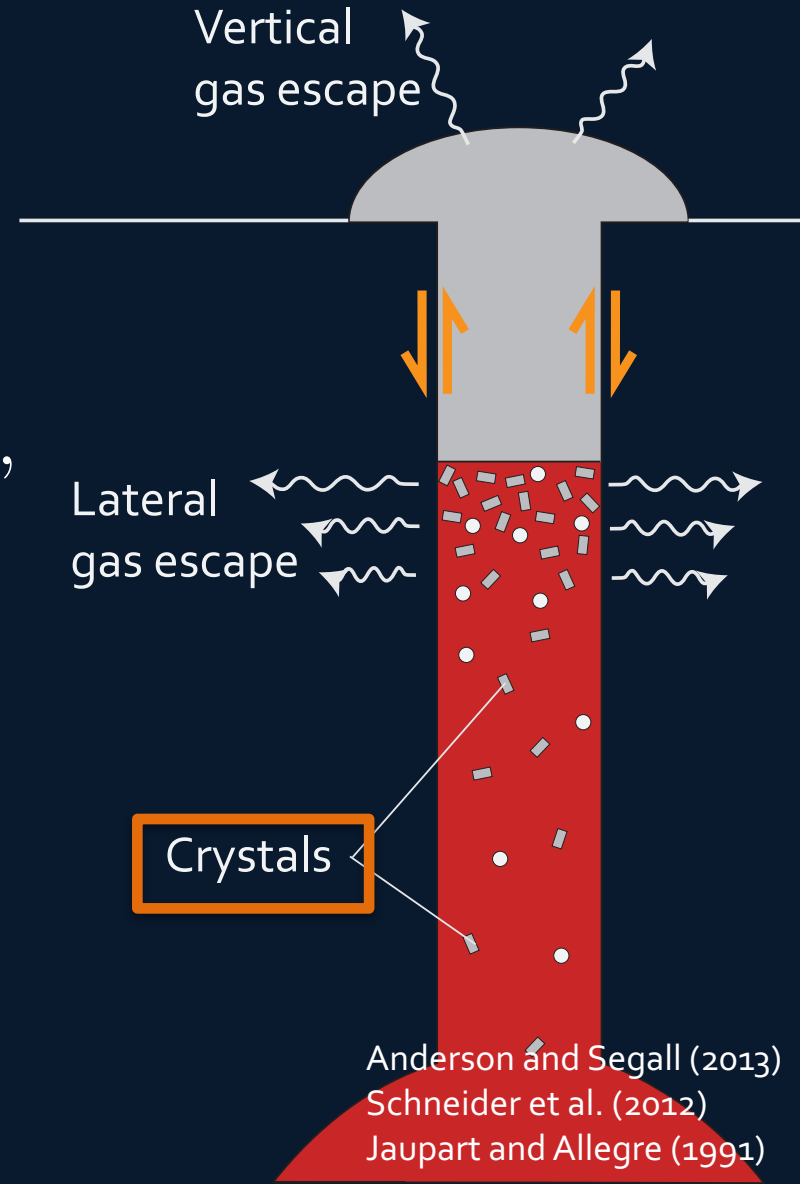
$$v = v_{\text{visc}} + v_{\text{fric}}$$

$$= \frac{\tau_R R}{4\eta} + 2v_{\text{ref}} \exp\left(\frac{\tau_R}{a\sigma} - \frac{f_0}{a}\right),$$

Unknown parameters

where

$$\tau_R = -\frac{R}{2} \left(\frac{\partial p}{\partial z} - \rho g \right)$$



Viscous flow to plug flow transition

$$v = v_{\text{visc}} + v_{\text{fric}}$$

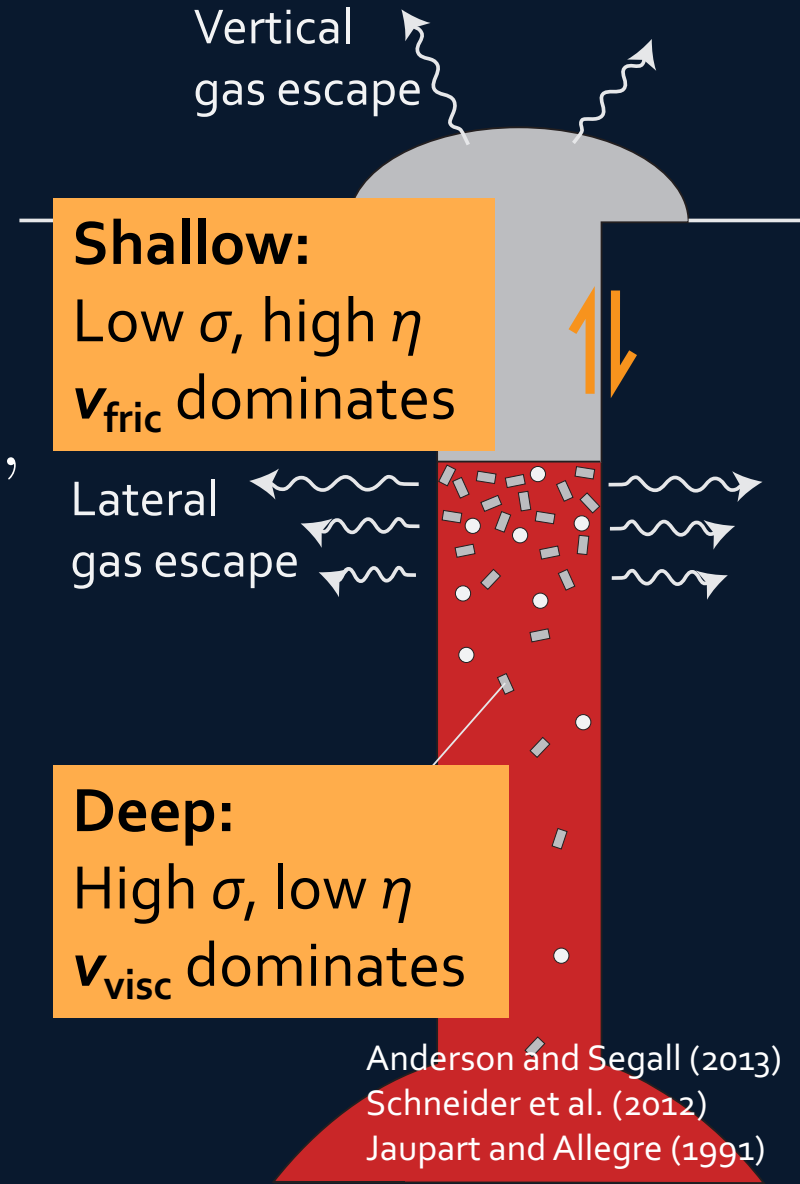
$$= \frac{\tau_R R}{4\eta} + 2v_{\text{ref}} \exp\left(\frac{\tau_R}{a\sigma} - \frac{f_0}{a}\right),$$

where

$$\tau_R = -\frac{R}{2} \left(\frac{\partial p}{\partial z} - \rho g \right)$$

Unknown parameters

f_0



Continuity Equations

Liquid + Solid:
$$\frac{\partial}{\partial t} (\rho_l \phi_l + \rho_s \phi_s) + \frac{\partial}{\partial z} [(\rho_l \phi_l + \rho_s \phi_s) v] = 0$$

Water:

$$\frac{\partial}{\partial t} \left(\chi_h^d \rho_l \phi_l + \frac{1}{1+\Gamma} \phi_g \rho_g \right) + \frac{\partial}{\partial z} \left[\left(\underbrace{\chi_h^d \rho_l \phi_l}_{\text{Dissolved}} + \underbrace{\frac{1}{1+\Gamma} \phi_g \rho_g}_{\text{Exsolved}} \right) v + \frac{1}{1+\Gamma} \phi_g \rho_g \underbrace{(v_g - v)}_{\text{Vertical Flow}} \right] = - \underbrace{\frac{2 \rho_g u_{lat}}{R (1+\Gamma)}}_{\text{Lateral Gas Loss}}$$

Analogous for CO₂:

$$u_{lat} = \phi_g \frac{k}{\eta_g} \frac{(p - p_{hyd})}{\lambda} \quad v_g - v = \frac{k}{\eta_g} \frac{\partial p}{\partial z}$$

Magma Permeability

$$k_m = k_c \phi^3, \phi > \phi_{gc}$$

Vertical gas escape through magma

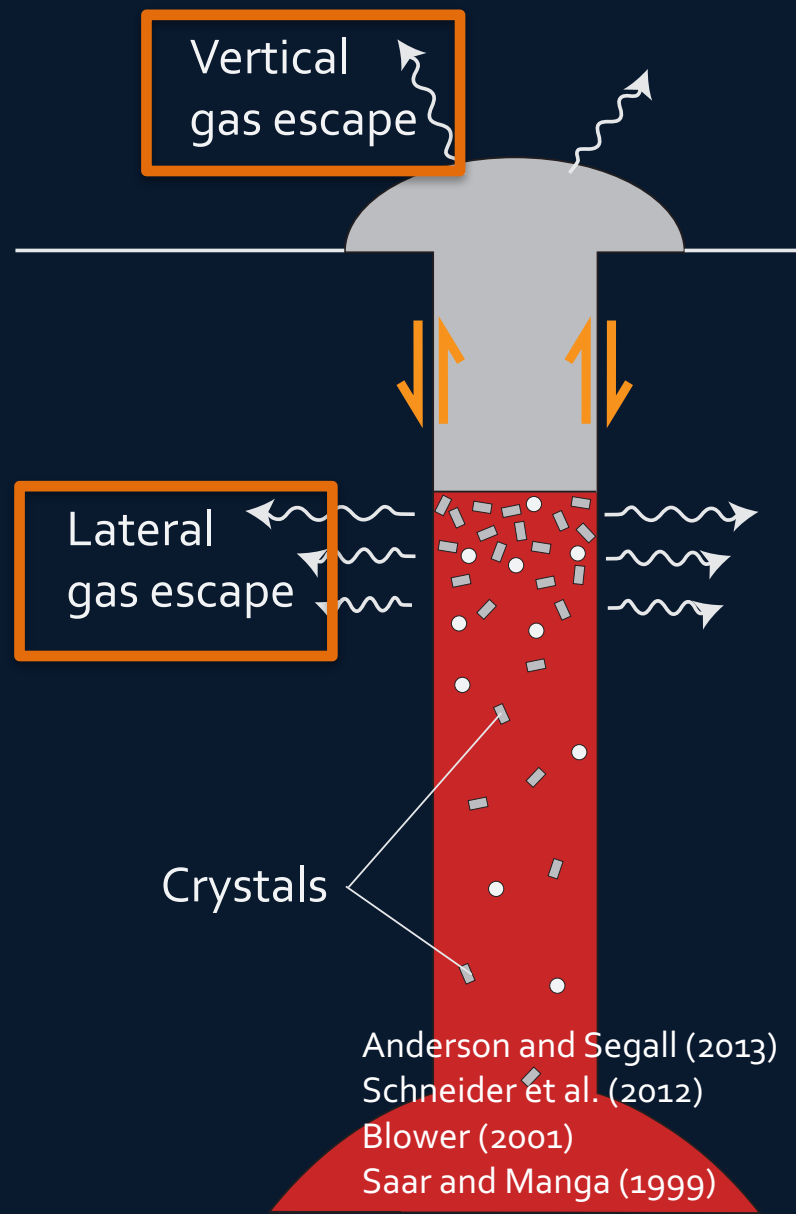
Percolation Theory:

$$k_m = k_c \phi^3, \phi > \phi_{ge}$$

Permeability
scaling
constant
(unknown)

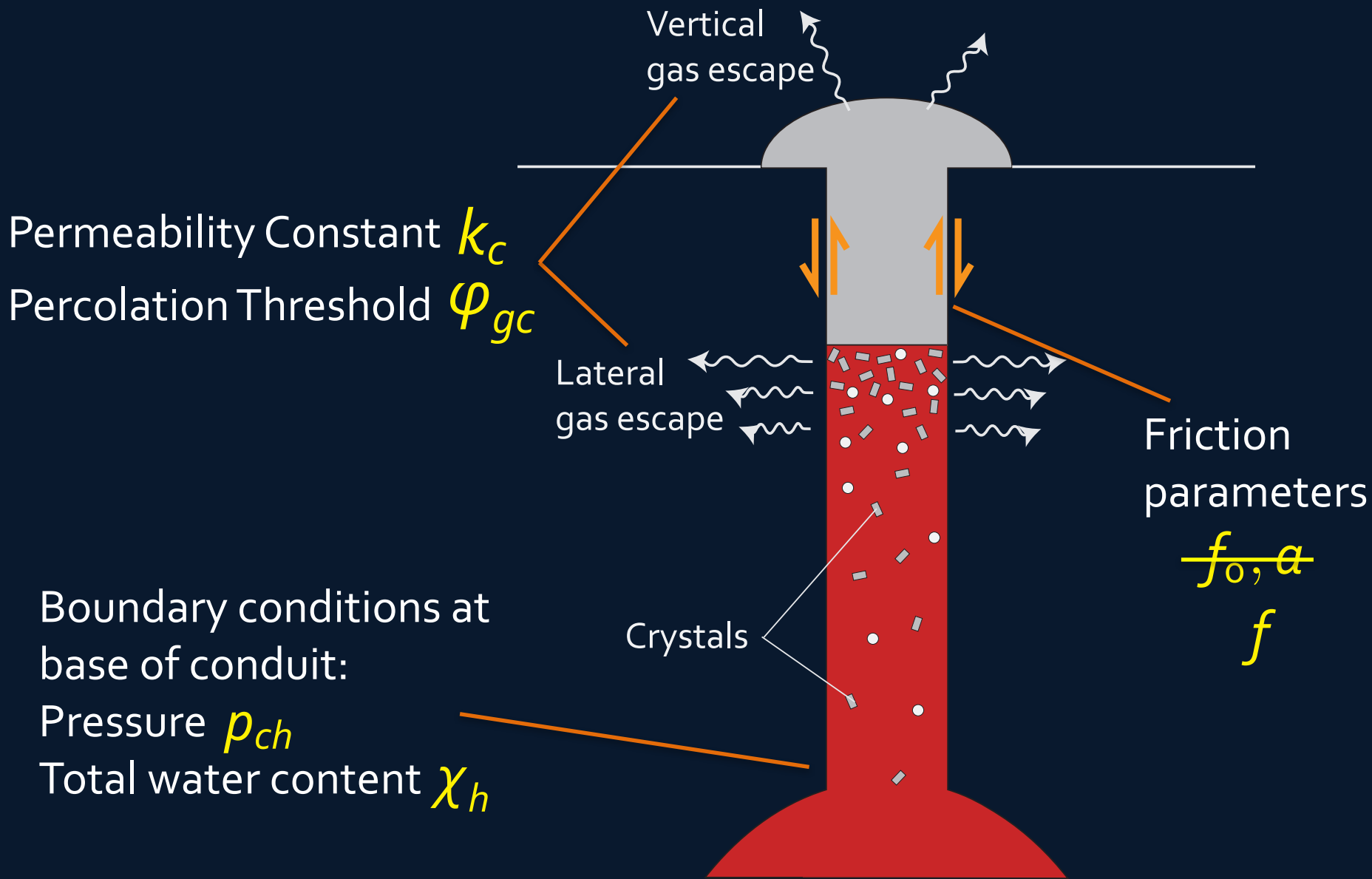
Porosity

Percolation
threshold
(unknown)

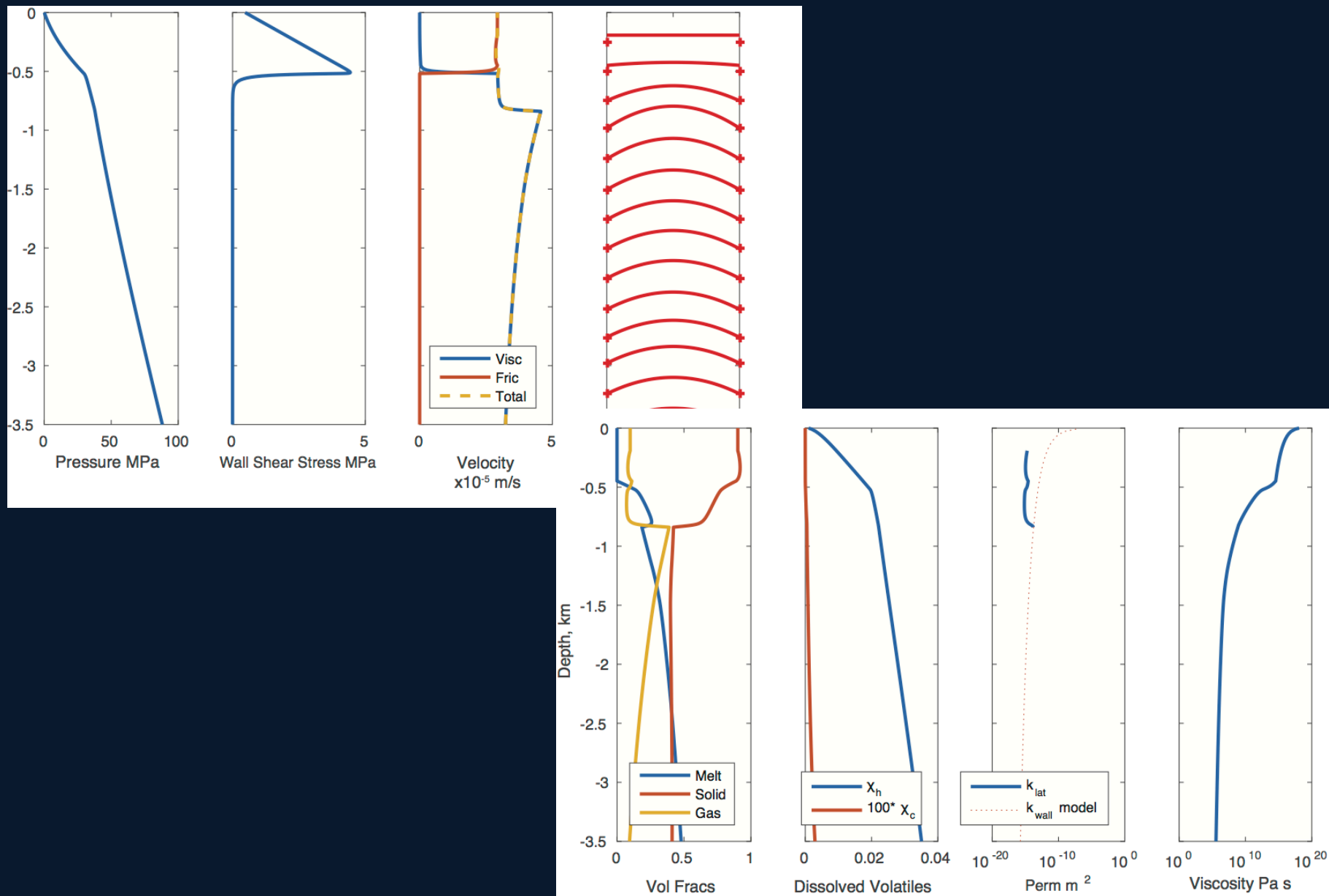


Anderson and Segall (2013)
Schneider et al. (2012)
Blower (2001)
Saar and Manga (1999)

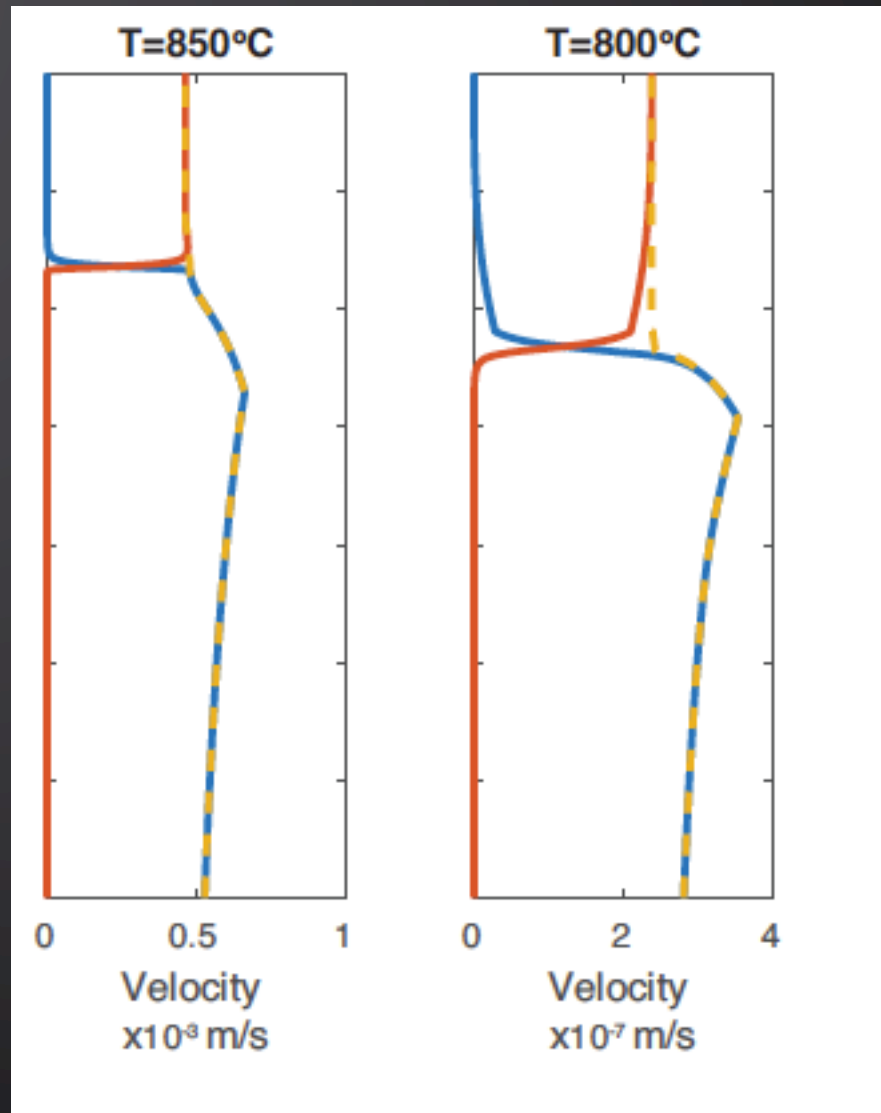
Summary of model parameters



Steady state solution

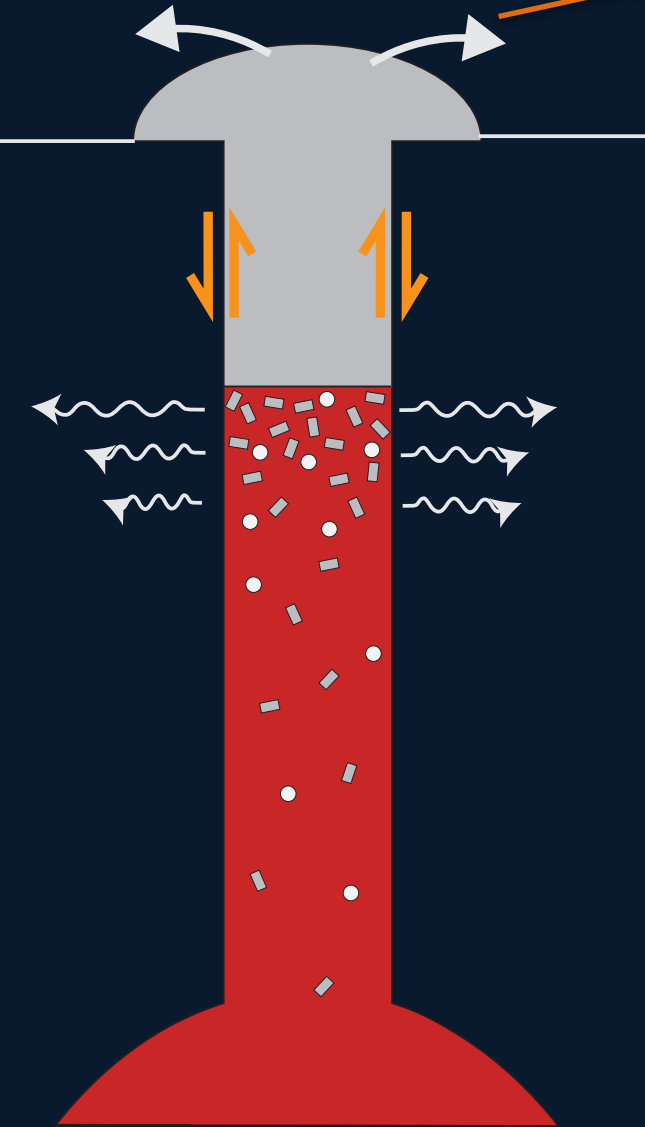


Temperature Dependence



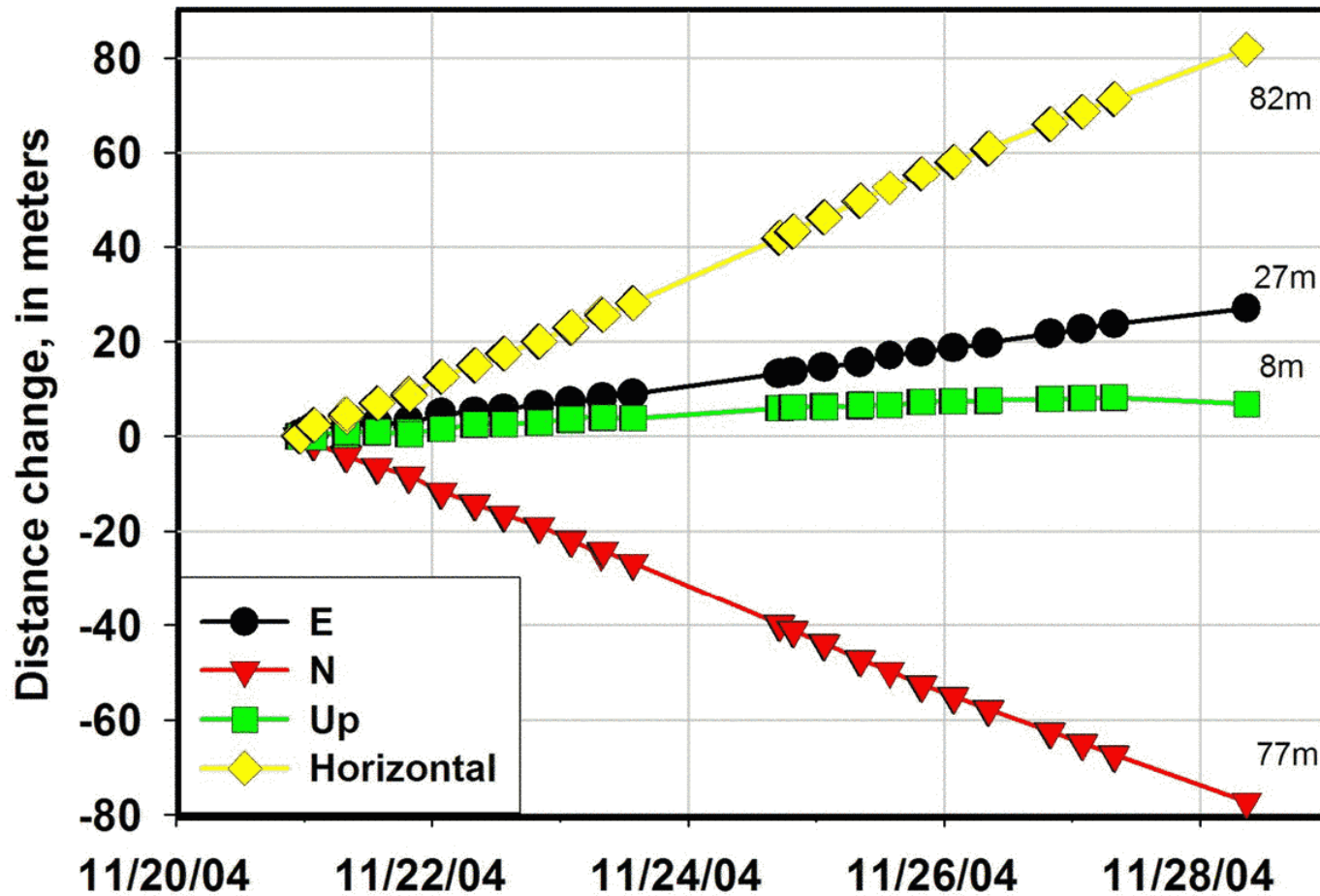
Quasi-steady state system: data

- 1) Exit velocity of plug
 $3 - 7 \times 10^{-5} \text{ m/s}$



Whaleback motion 11/21/04 – 11/29/04

GPS spider station ELEA on 2004 Dome

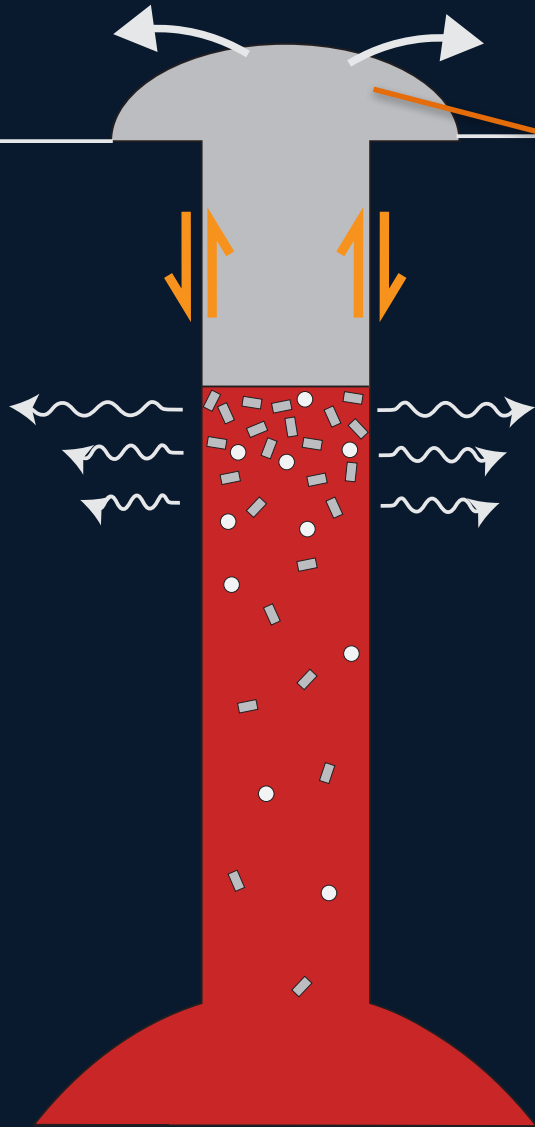


Quasi-steady state system: data

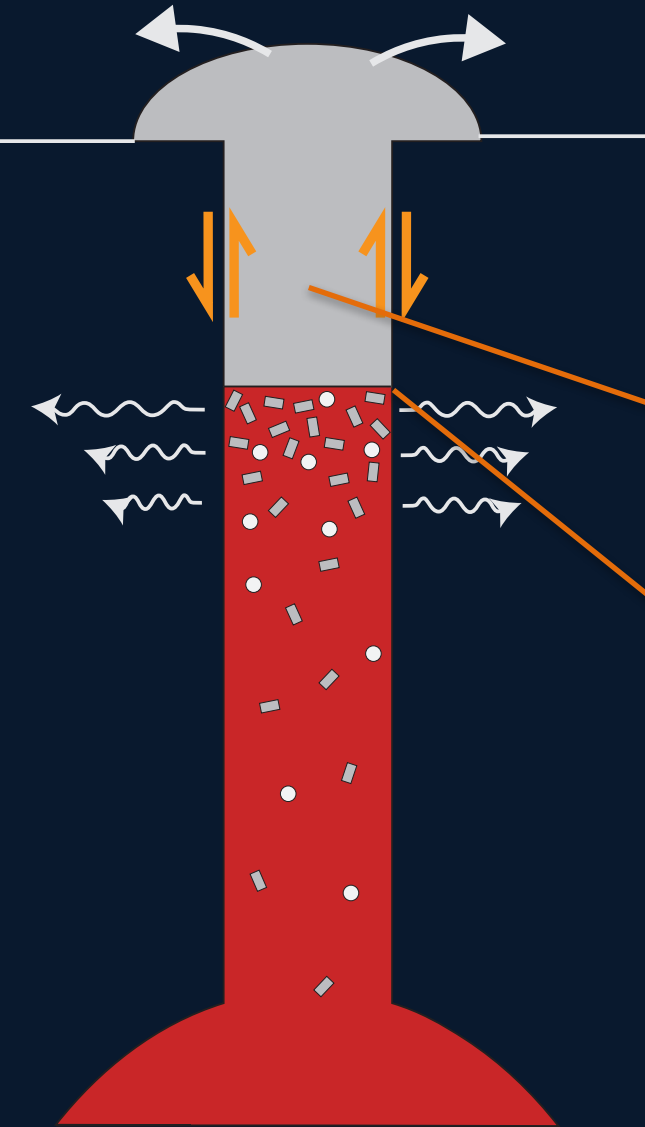
1) Exit velocity of plug
 $3 - 7 \times 10^{-5} \text{ m/s}$

2) Dome rock porosity
 $5 - 10\%$

Thornber et al. (2008), Cashman et al. (2008)



Quasi-steady state system: Data



1) Exit velocity of plug
 $3 - 7 \times 10^{-5} \text{ m/s}$

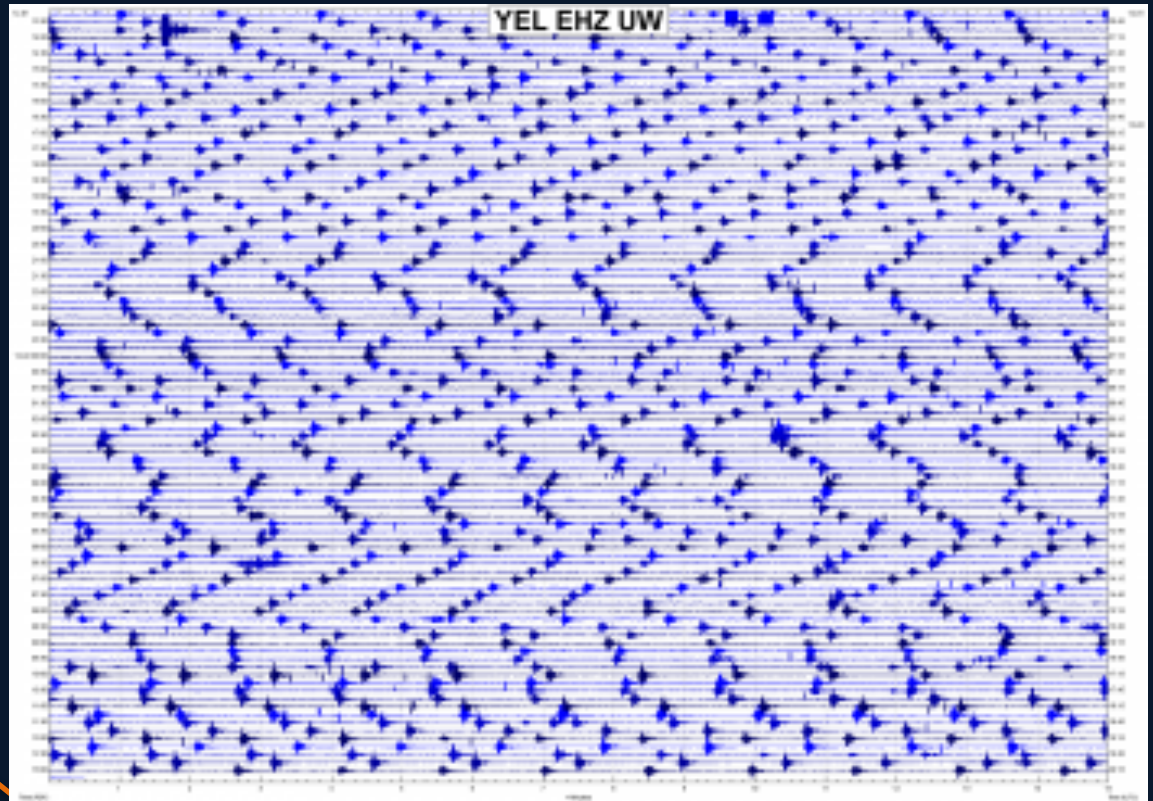
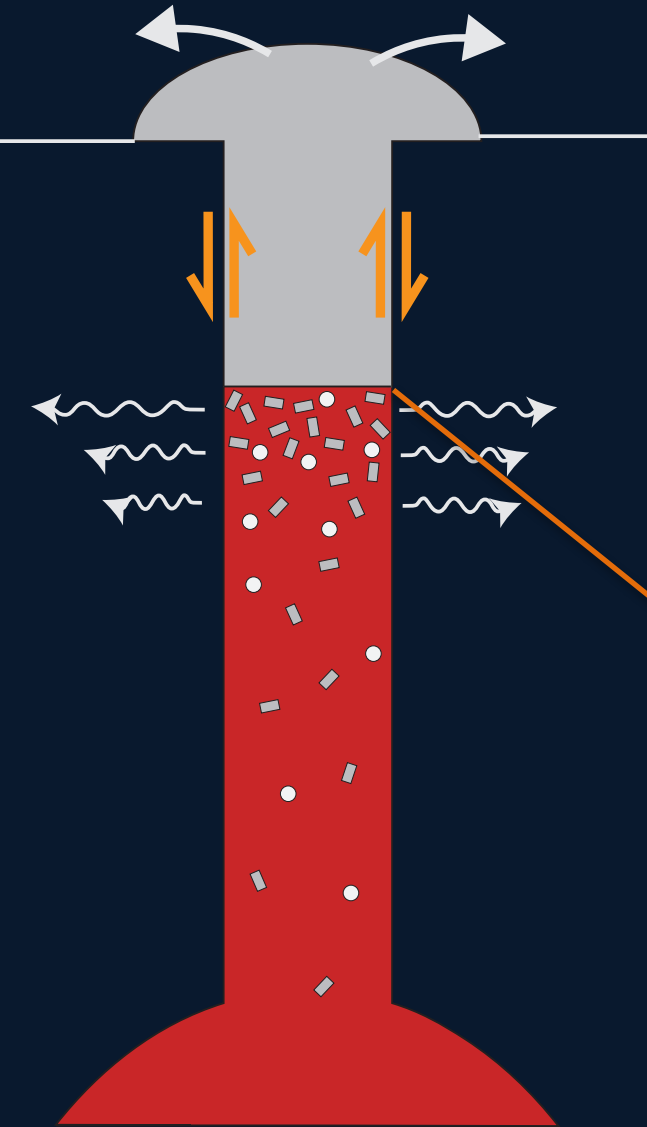
2) Dome rock porosity
 $5 - 10\%$

3) Crystallization depth
 $0.5 - 1 \text{ km}$

4) Plug depth
 $0.5 - 1 \text{ km}$

Vallance et al. (2008), Schilling et al. (2008)
Thornber et al. (2008), Cashman et al. (2008)
Pallister et al. (2008)
Iverson et al. (2006), Moore et al. (2008)

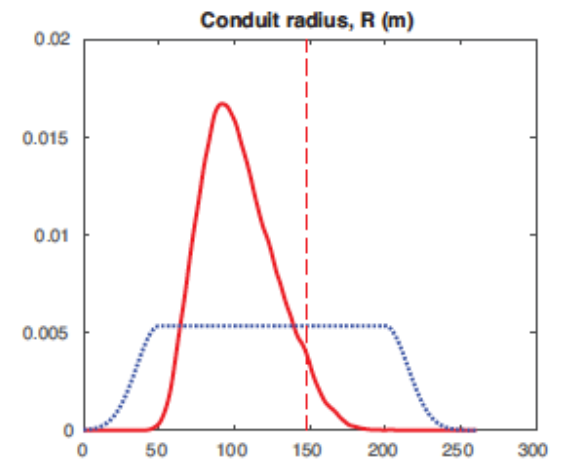
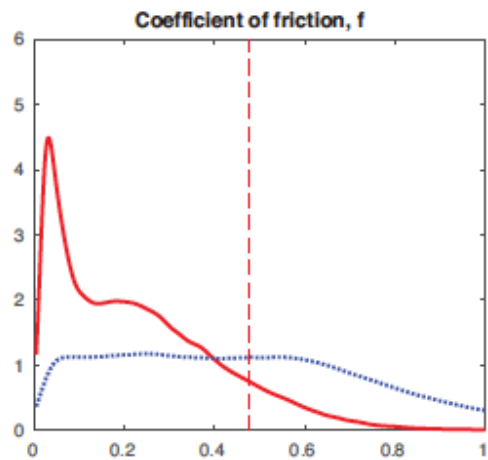
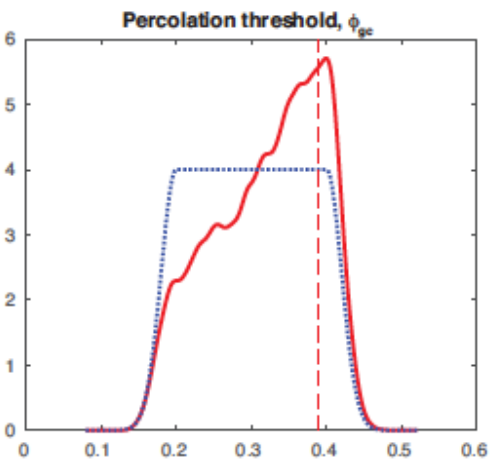
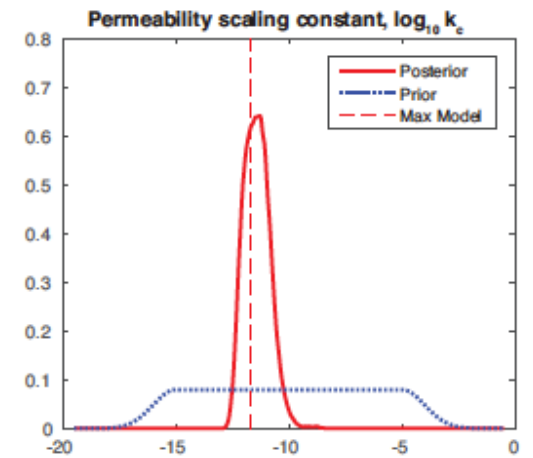
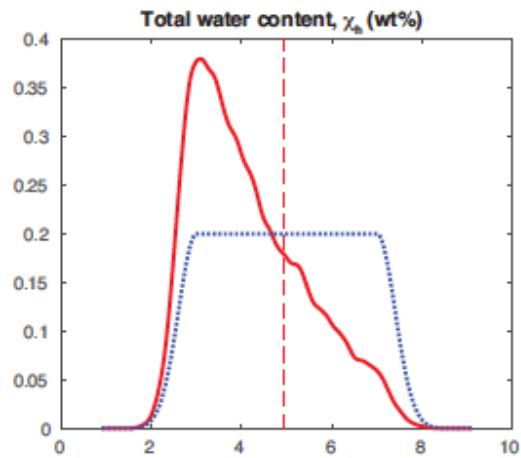
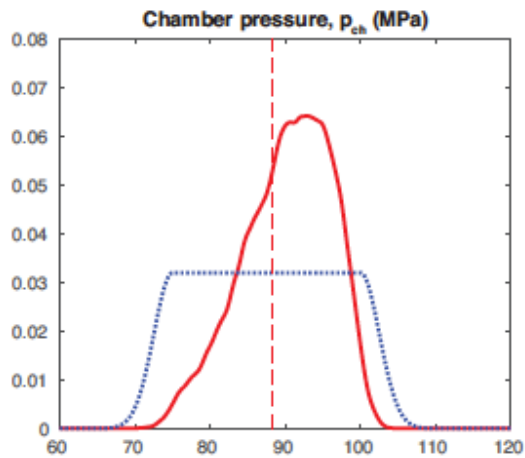
Quasi-steady state system: Data



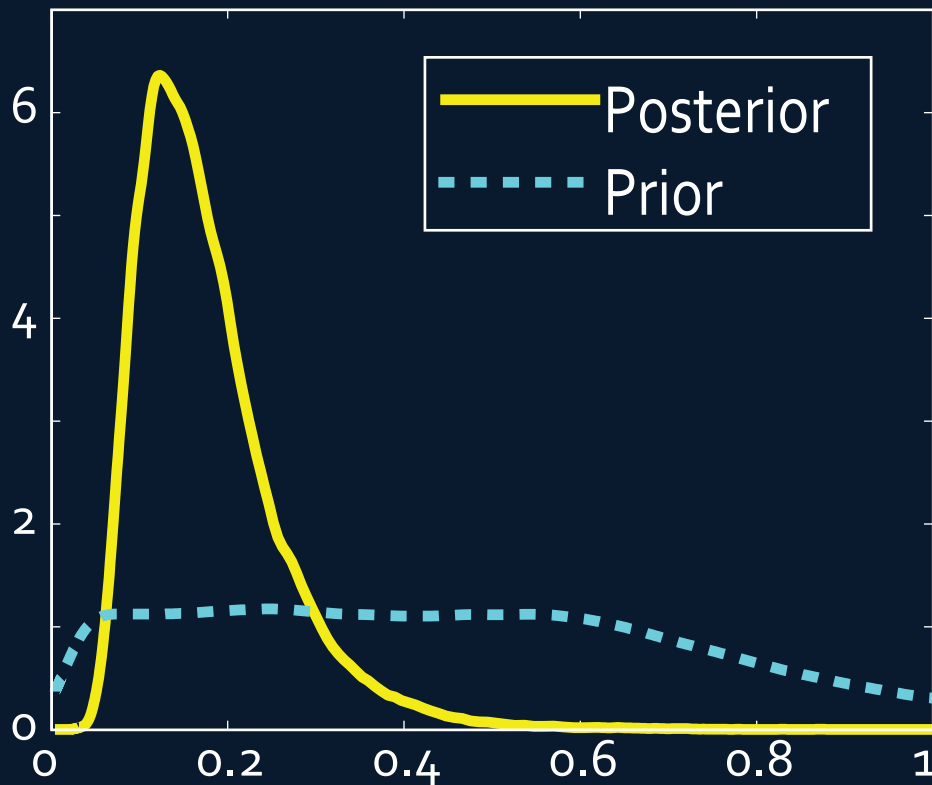
4) Plug depth
0.5 – 1 km

Vallance et al. (2008), Schilling et al. (2008)
Thorner et al. (2008), Cashman et al. (2008)
Pallister et al. (2008)
Iverson et al. (2006), Moore et al. (2008)

Results: Posterior PDFs



Low friction along conduit wall



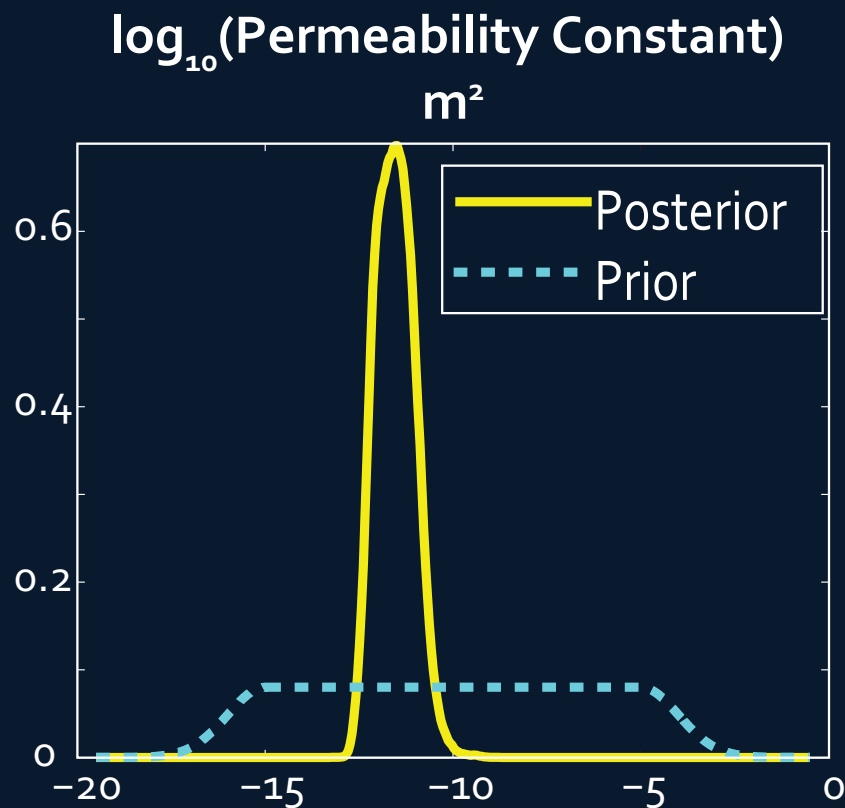
Experiments from
Moore et al. (2008)

- Most accepted models $f < 0.3$

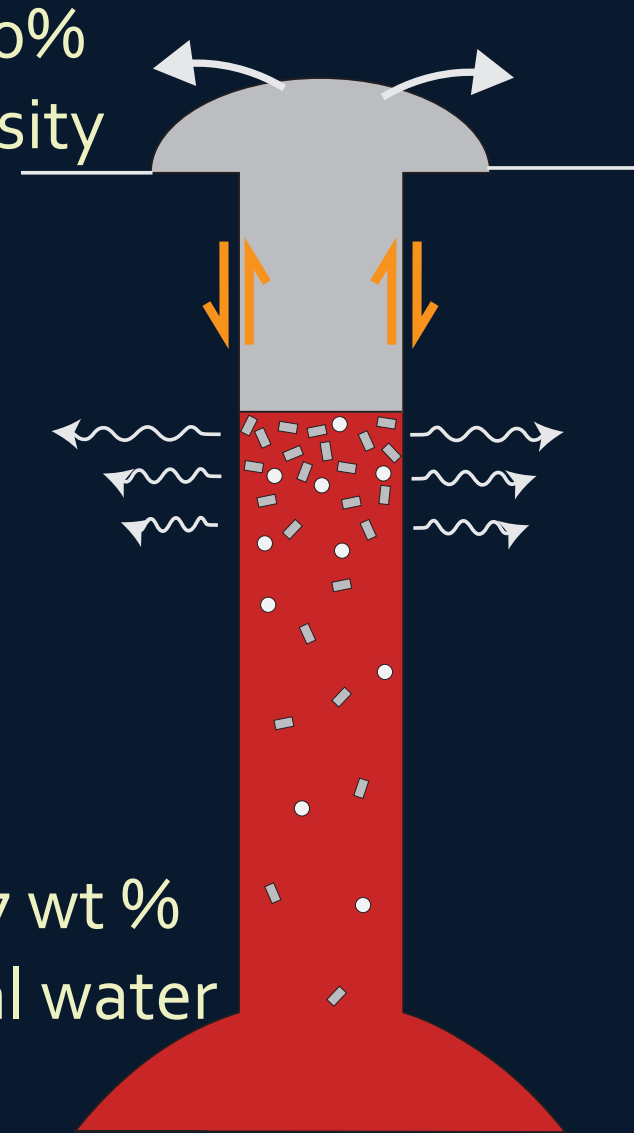
$$f = \frac{\tau_R}{\sigma - p} = \frac{\tau_R}{\sigma_{\text{eff}}}$$

- Possible explanations
 - High pore pressure
 - Reduced normal stress
 - Bias in viscosity

Well-constrained magma permeability constant



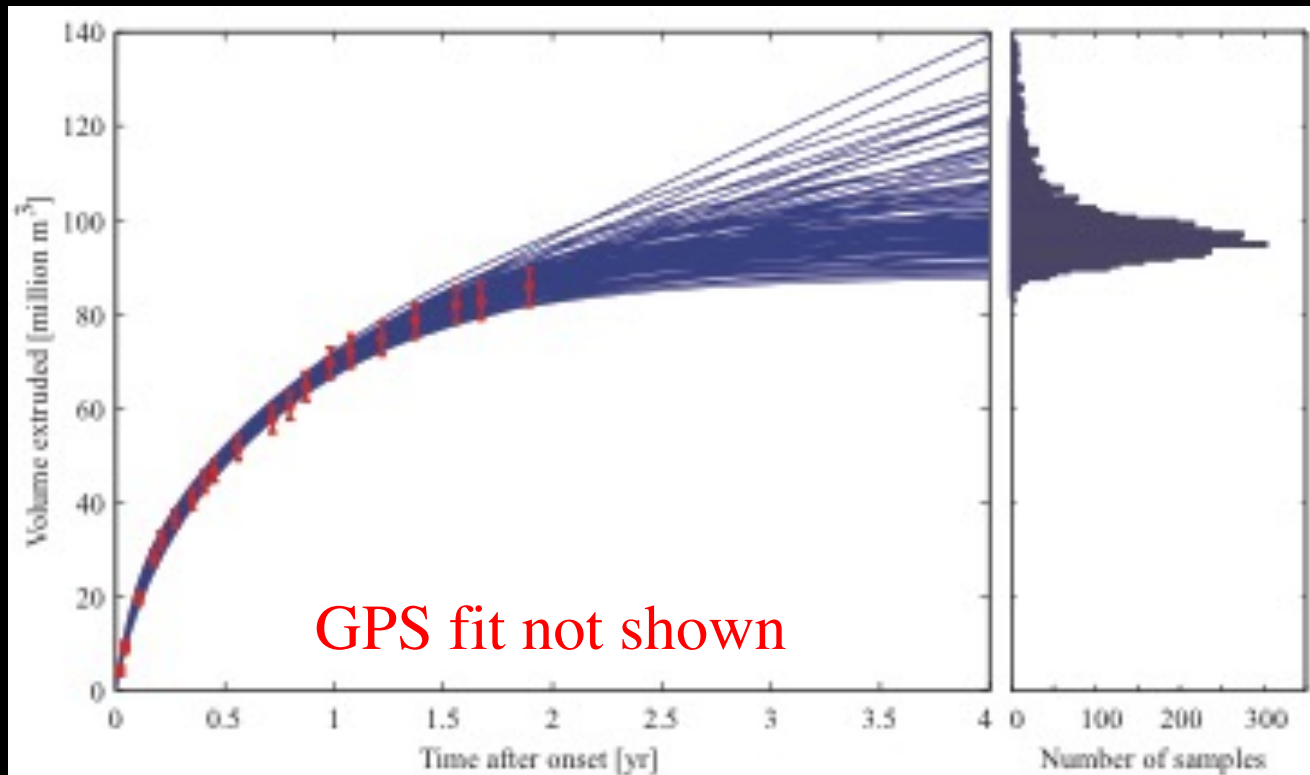
5 – 10% porosity



Can We Forecast the Size and Duration of an Eruption?

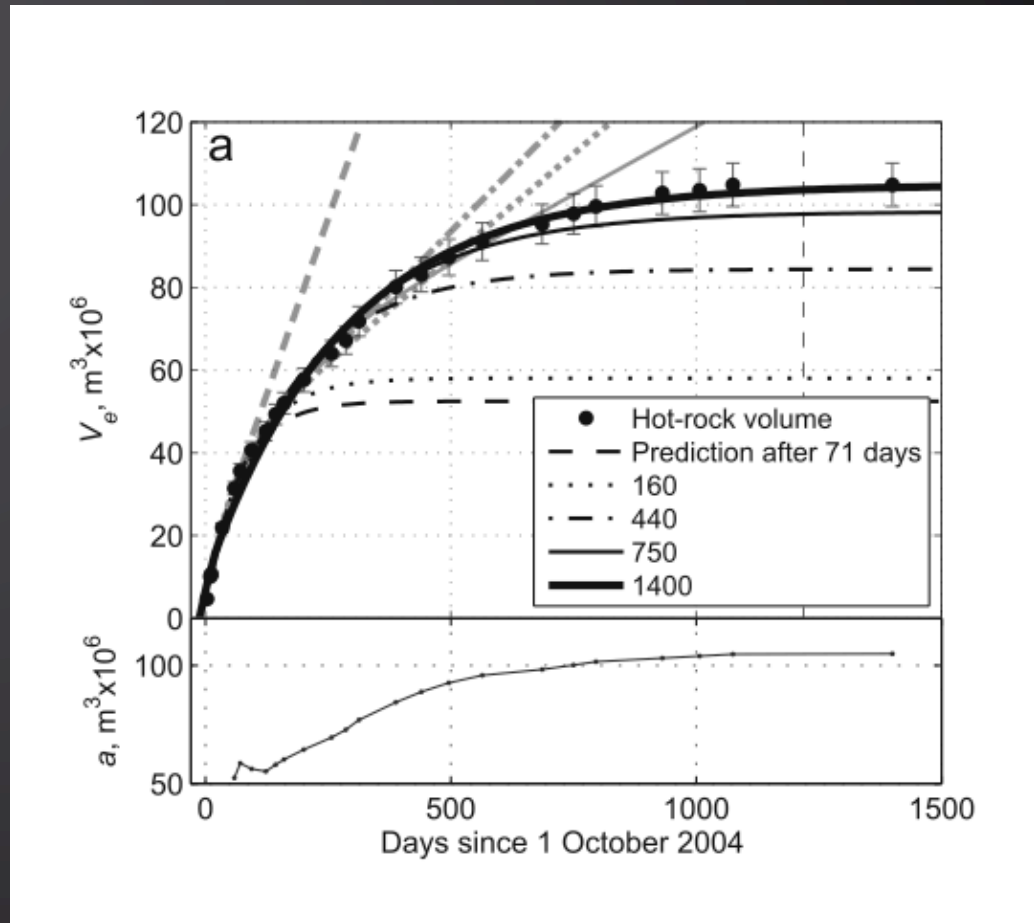


Monte Carlo Forecasting



- Forecast based on knowledge of the system and all existing data.
- Yields probabilistic forecast including uncertainties in the underlying parameters (not “epistemic” uncertainty).
- Uncertainty increases with time.

“Real Time” Forecasting at Mount St Helens



$$\frac{V_{ex}(t)}{V_0 \bar{\beta} (p_0 - \rho g L)} = \left(\frac{\alpha}{\Omega + \alpha} \right) \left[\left(\frac{\alpha}{\Omega + \alpha} \right) (1 - e^{-t/t_c}) + \frac{t}{t_c} \left(\frac{\Omega}{\Omega + \alpha} \right) \right]$$

Conclusions

- Combining geodetic data and physics-based models constrains more properties of magmatic systems than geodetic data alone.
- It is possible to include other data types (gas emission, gravity,) to better constrain systems.
- *May* be possible to use in forecasting mode, but open question.
- Model components need to be well understood.