Monsoons in NICAM and challenges ahead

Tomoe Nasuno

Japan Agency for Marine-Earth Science and Technology

ICTP-IITM-COLA Targeted Training Activity (TTA) "Towards Improved Monsoon Simulations" June 13 - 17, 2016 Miramare, Trieste, Italy

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3. Challenges ahead

NICAM: Nonhydrostatic Icosahedal Atmospheric Model

- Development since 2000
 Tomita and Satoh (2005, Fluid Dyn. Res.)
 Satoh et al. (2008, J. Comp. Phys.)
- First global dx=3.5km run in 2004

 using the Earth Simulator (JAMSTEC)
 Tomita et al.(2005, *Geophys. Res. Lett.*)
 Miura et al.(2007, Science)
- K computer era (10PF; Kobe,Riken,2012)
 Toward higher resolution: dx=870m (Miyamoto et al. 2013)

 Multi ensemble simulations
 (Miyakawa et al. 2014 Nat. Comm; Nakano et al. 2015, Geophys. Res. Lett.)

Multi decadal simulations (Kodama et al. 2015, JMSJ)



http://nicam.jp/hiki/



NICAM outcomes: 10-year history and beyond

- Good points of NICAM
 - Realistic meso-scale circulations, e.g. diurnal cycle
 - Multiscale structure of cloud systems
 - Intra-seasonal oscillation: MJO, BSISO (boreal summer ISO)
 - Tropical cyclones
 - Cloud properties with cloud microphysics
 - Collaboration with satellite observation (evaluation, improvements, & assimilation)



Progress in Earth

Science

9

and Planetary

Overview paper:

Satoh, M., Tomita, H., Yashiro, H., Miura, H., Kodama, C., Seiki, T., Noda, A. T., Yamada, Y., Goto, D., Sawada, M., Miyoshi, T., Niwa, Y., Hara, M., Ohno, T., Iga, S., Arakawa, T., Inoue, T., Kubokawa, H. (2014) The Nonhydrostatic Icosahedral Atmospheric Model: Description and Development.

Progress in Earth and Planetary Science, **1**, 18. http://dx.doi.org/10.1186/s4064501400181

Nonhydrostatic Icosahedarl Atmospheric Model

Satoh et al. (2014)

Dynamics

| governing equations | Fully compressible non-hydrostatic system | | | | |
|--|---|--|--|--|--|
| spatial discretization horizontal grid configuration vertical grid configuration topography | Finite Volume MethodIcosahedral grid (Tomita et al. 2001, 2002)Lorenz gridTerrain-following coordinate | | | | |
| conservation | Total mass, total energy Satoh (2002, 2003) | | | | |
| temporal scheme | Slow mode — explicit scheme (RK2, RK3) Fast mode — Horizontal Explicit Vertical Implicit scheme | | | | |
| ■ Physics | • | | | | |
| radiation | MSTRNX / MSTRNX-AR5 (Sekiguchi and Nakajima, 2008) | | | | |
| cloud physics | Grabowski(1998); NSW6(Tomita 2008);NDW6(Seiki et al. 2013) | | | | |
| shallow clouds boundary layer | MY level 2 (Mellor and Yamada 1982; Noda et al. 2010) MYNN level 2.5 or 3 (Nakanishi and Niino 2006) | | | | |

surface flux Louis(1979), Uno et al.(1995)

surface processes SST specified & bucket / slab ocean & MATSIRO

NICAM 870m-mesh simulation

- Miyamoto, Y., Kajikawa, Y., Yoshida, R., Yamaura, T., Yashiro, H., and Tomita, H., 2013: Deep moist atmospheric convection in a subkilometer global simulation. Geophys. Res. Lett., 40, 4922-4926. DOI:10.1002/grl.50944.
- Miyamoto, Y., R. Yoshida, T. Yamaura, H. Yashiro, H. Tomita and Y. Kajikawa, 2015: Does convection vary in different cloudy disturbances? Atmospheric Science Letters, 16, 305-309, DOI:10.1002/asl2.558.
- Kajikawa, Y., Miyamoto, Y., Yoshida, R., Yamaura, T., Yashiro, H., Tomita, H., 2016: Resolution dependence of deep convections in a global simulation from over 10-kilometer to sub-kilometer grid spacing. Progress in Earth and Planetary Science, accepted.

OLR (6UTC, 25 Aug. 2012)



NICAM 870m-mesh simulation Miyamoto et al. (2013,GRL) using the K computer



dx=3.5 m





Global View: 0600 UTC, 25, Aug. 2012







SNIPER Project

Composite of convection (vertical velocity)



 $\Delta x \ge 3.5$ km:

- Convection is represented at <u>1</u> grid
- Little dependence on resolution
- $\Delta x \leq 1.7$ km:
 - Convection is represented at <u>multiple grids</u>
 - Intensify w/ resolution

*transform the coordinate into the cylindrical around the core grid mean of all the detected convection

symmetric around the x axis X axis is normalized by resolution

Miyamoto et al. (2013)

2. Monsoon simulations in NICAM Indo-Asian Summer Monsoon

- Significant seasonal cycle in broad tropical and subtropical domain, affecting world weather and climate (e.g., teleconnection, air-sea interaction)
- Local onset period: late May-early June (south India, SCS), late July (WNP, end of Baiu), Climatologically phase-locked meridional migration (Murakami and Matsumoto 1994; Ueda et al. 1995)
- Multiscale nature: TC activity is closely related to the convective activity associated with monsoon (e.g.,Ueda et al. 1995). Monsoon circulation is a basic background of TC genesis in WNP (Holland 1995).
- Decadal scale modulation in SCS/WNP monsoon onset date associated with SST variability (Kajikawa and Wang 2012; Tomita et al. 2013).



Monsoon simulations in NICAM Indian Summer Monsoon

- Kajikawa, Y., Yamaura, T., Tomita, H., Satoh, M., 2015: Impact of tropical disturbance on the Indian summer monsoon onset simulated by a global cloud-system-resolving model. SOLA, 11, 80-84, doi:10.2151/sola.2015-020.
- <u>Kinter III, J. L., et al., 2013</u>: Revolutionizing Climate Modeling -Project Athena: A Multi-Institutional, International Collaboration. Bull. Am. Meteorol. Soc., 94, 231-245. <u>http://journals.ametsoc.org/doi/abs/10.1175/BAMS-D-11-</u> 00043.1
- <u>Satoh, M, et al., 2012</u>: The Intra-Seasonal Oscillation and its control of tropical cyclones simulated by high-resolution global atmospheric models, Clim. Dyn., 39.2185-2206, DOI 10.1007/s00382-011-1235-6.
- <u>Taniguchi, H., W. Yanase, M. Satoh, 2010</u>: Ensemble simulation of cyclone Nargis by a Global Cloud-system-resolving Model -modulation of cyclogenesis by the Madden-Julian Oscillation. J. Meteor. Soc. Japan, 88, 571-591.

The Athena Project Kinter et al.(2013)



Collaborating Groups

- **COLA** Center for Ocean-Land-Atmosphere Studies, USA
- **ECMWF** European Center for Medium-range Weather Forecasts, UK
- **JAMSTEC** Japan Agency for Marine-Earth Science and Technology, Research Institute for Global Change, Japan
- University of Tokyo, Japan
- NICS National Institute for Computational Sciences, USA
- Cray Inc.

Codes

- NICAM: Nonhydrostatic Icosahedral Atmospheric Model
- IFS: ECMWF Integrated Forecast System

Super-computers

- Athena: Cray XT4 4512 quad-core Opteron nodes (18048)
- Kraken: Cray XT5 8256 dual hex-core Opteron nodes (99072)

Athena Experiments JJA in 2001-2009 (initialized each year)

| | Model/Exp. | Resolution | # Cases | Period | Notes | |
|---|------------------------------|-----------------------------------|---------|----------|---|--------|
| < | NICAM / Hindcasts | 7 km | 8 | 103 days | 21 May - 30 Aug 2001 - 2009 | \sum |
| | | | | | | |
| | IFS / Hindcasts | 125 km 39 km 16 km | 48 | 395 days | 1 Nov - 30 Nov (following year) 1960 - 2007 | |
| | IFS / Hindcasts | 10 km | 20 | | 1 Nov - 30 Nov (following year) 1989 - 2007 | |
| < | IFS / Hindcasts | 125 km 39 km 16 km 10 km | 9 | 103 days | 21 May - 30 Aug 2001 - 2009 NICAM analogs | \sum |
| | IFS / Summer Ensembles | 39 km 16 km | 6 | 132 days | 21 May - 30 Sep selected years | |
| | IFS / Winter Ensembles | 39 km 16 km | 6 | 151 days | 1 Nov - 31 Mar selected years | |
| | IFS / AMIP | 39 km 16 km | 1 | 47 years | 1961 - 2007 | |
| | IFS / Time Slice | 39 km 16 km | 1 | 47 years | 2071 - 2117 | |

http://wxmaps.org/athena/home/

JJA Precipitation (anomaly from GPCP)

IFS

Dirmeyer et al (2012) Clm. Dyn.



NICAM

GPCP



Fig. 1 The 8-year averaged monthly mean zonal wind at 850 hPa in NICAM (top) and IFS (middle) simulations in comparison with ERA–Interim (bottom).



Satoh et al. (2012)



Fig.3 Time series of the Indian Monsoon Index (Wang et al. 2001) for NICAM and IFS simulations and ERA–Interim data. (a) The 8-year average and (b)–(i) anomaly from the average are plotted. 5-day running mean is operated.

Satoh et al. (2012)



Fig.2 Time-latitude sections of the 8-year 60-90E averaged precipitation (color) and zonal wind at 850 hPa (contour lines) in NICAM (top) and IFS (middle) simulations in comparison with TRMM–3B42 and ERA–Interim and (bottom). Contour intervals for Zonal wind is 4 m s-1 (solid: positive, broken: negative). Zero contour lines are omitted.



Fig.4 Time-latitude sections of anomalous 60-90E average surface precipitation (color) and zonal wind at 850 hPa (contour lines in the initial 52 days of NICAM (top) and IFS (middle) simulations in comparison with TRMM–3B42 and ERA–Interim data (bottom). The anomalies from the 8-year average (Fig. 2) are plotted. Contour intervals for zonal wind are 2 m s⁻¹(solid: positive, broken: negative). Zero contour lines are omitted.



Fig.4 (continue)



Satoh et al. (2012)

Fig. 5 Pentad-mean surface precipitation and wind vectors at 850 hPa for (a) 15–19 and (b) 25–29 June 2006 for NICAM (top) and IFS (middle) simulations in comparison with TRMM-3B42 and ERA-Interim data (bottom).









Summary (JJA ensemble simulations; Satoh et al. 2012)

- In the first month of simulation, both models capture the intra-seasonal oscillatory behavior of the Indian monsoon similar to the observed boreal summer ISO in approximately half of the 8-year samples.
- The IFS simulates the NW–SE-oriented rainband and the westerly location better, while NICAM marginally reproduces mesoscale organized convective systems and better simulates the northward migration of the westerly peak and precipitation, particularly in 2006.

Impact of tropical disturbance on the Indian summer monsoon onset in 2012 (Kajikawa et al. 2015)

Indian summer monsoon indices



Kajikawa et al. (2015)



Mean of Exp. with IC at May 15-19



Kajikawa et al. (2015)

Kajikawa et al. (2015)



- Summary (onset of ISM in 2012; Kajikawa et al. 2015)
- Focuses are on the effect of tropical disturbances on ISM onset and considered the potential extension of onset predictability
- A series of 30-day simulations initialized on10 May-10 June show a skill of two weeks predictability of ISM onset.
- The ISM onset was accompanied by northward-migrating tropical disturbances over the Bay of Bengal and the Arabian Sea, which originated in the equatorial Indian Ocean.
- The result suggest that a better representation of tropical disturbances enhances the potential to extend the predictability of the transition phase of the Asian summer monsoon.

Monsoon simulations in NICAM East Asian Summer Monsoon

- Kodama, C., Yamada, Y., Noda, A. T., Kikuchi, K., Kajikawa, Y., Nasuno, T., Tomita, T., Yamaura, T., Takahashi, T. G., Hara, M., Kawatani, Y., Satoh, M., Sugi, M.2015: A 20-year climatology of a NICAM AMIP-type simulation. J. Meteor. Soc. Japan, 93, 393-424, doi:10.2151/jmsj.2015-024.
- Yamaura, T., Kajikawa, Y., Tomita, H., Satoh, M., 2013: Possible impact of a tropical cyclone on the northward migration of the baiu frontal zone. SOLA, 9, 89-93. <u>http://dx.doi.org/10.2151/sola.2013-020</u>
- Oouchi, K.,A. T. Noda, M. Satoh, B. Wang, S.-P. Xie, H.G. Takahashi, T. Yasunari (2009): Asian summer monsoon simulated by a global cloud-system-resolving model: Diurnal to intra-seasonal variability. Geophys. Res. Lett.,36, L11815, doi:10.1029/2009GL038271.

NICAM AMIP-type simulations

Kodama et al. (2015)

- 14km horizontal mesh and 38 vertical levels up to 40 km.
- 1-moment 6-category bulk cloud microphysics (Tomita 2008).
 - No cumulus convection parameterization
 - parameters tuned by several seasonal-scale experiments
- AMIP configurations except for
 - slab ocean model (D=15m & τ=7days) with SST nudging and fixed sea ice
- CNTL run: 1978.06-2009.12
 - monthly mean AMIP2 SST/SSI.
- FUTURE run: 2074.06-2105.12 (A1B scenario)
 - CMIP3 model ensemble dSST = SST(2075-2099) SST(1979-2003) including trend is added to AMIP2 SST. For sea ice, areal change is considered following Mizuta et al. [2008].

precipitation: SLAB vs. fixed SST



Precipitation rate



0



Fig. 1. A: Difference between OLR(max) and OLR(min). Intervals are for 20 Wm⁻²; hatching denotes OLR(min) greater than 240 Wm⁻², while dark (light) shading indicates regions of DD greater than 60 (40) Wm⁻². B: The domains of the three monsoon systems SEAM, WNPM and NAIM, as well as two extratropical wet-climate regimes of the TIBU and BAIU. Refer to the text for further information.



precipitation JJAS 20-yr Climatology



Climatology of Asian monsoon U850(1979-1998)



OLR annual cycle of 20-yr Climatology



Kodama et al. (2015, JMSJ)





Kodama et al. (2015, JMSJ)



Fig. 17 Climatological mean time series of maximum $-\partial \langle \theta e \rangle / \partial y$ latitude averaged between 125° and 145° E (stepwise solid line; left axis), $\langle \theta e \rangle$ averaged from 125 to 145° E, and 30 to 40° N (smoothed dotted line; right axis), and $\langle \theta e \rangle$ averaged from 125 to 145° E, and 20 to 30° N (smoothed dashed line; right axis). JRA-25 results are shown in (a), and NICAM results in (b). Left axis shows maximum $-\partial \langle \theta e \rangle / \partial y$ latitude in degrees north, and the 30° N and 40° N latitudes are indicated as horizontal solid lines. Right axis shows $\langle \theta e \rangle$ in K, and the 335 K $\langle \theta e \rangle$ is indicated as horizontal dashed line.



- Summary (AMIP-type simulation; Kodama et al. 2015)
- The 20-yr climatology of the seasonal march of Asian monsoon is reasonably reproduced (much better than the 8-year JJA simulations).
- Biases: northward displacement of westerly axis and subtropical high, which affects monsoon subsystems (e.g., the earlier and shorter Baiu).
- The interannual variability of WNP monsoon circulation (1979-2008) was simulated at correlation coefficient of 0.59 with JRA25). The performance corresponds to that in convection.

3. Challenges ahead

Multi-scale interactions:

- monsoon onset, active/break phase transition
- ISO, Tropical Cyclones

Inter-annual variability:

- ENSO
- Inter-decadal variabilities (PDO, NAO)

Global monsoons

- ✓ Model development (air-sea interactions, land surface processes)
- ✓ Metrics (multi-scale interactions)
- ✓ Model biases (evaluation, understanding)





Figure 3. (a) Observed and (b) simulated correlation coefficients between the June–August SST and precipitation anomalies (the color shadings). The contours denote the climatological June–August mean rainfall rate (in units of mm day⁻¹). The observed correlations were computed using 20 years of data (1982–2001) derived from CMAP rainfall and Reynolds SST. The simulated results were made by 5 AGCM's multi-model ensemble simulation.

1984

NICAM

SST(obs) ANOM JJAS 1984



"OLR anom





NICAM

SST(obs) ANOM JJAS 1985



OLR anom



60N



NICAM

SST(obs) ANOM JJAS 1996



OLR anom





NICAM

-0.2

-0.4

-0.8

-0.6

0.2

0.6

0.B

0.4

SST(obs) ANOM JJAS 1997







