

# **Influence of midlatitude disturbances on the MJO**

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Advanced School and Workshop on Tropical-Extratropical  
Interactions on Intra-seasonal time scales  
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# Outlines

- Introduction
- Tropical processes and MJO
- Midlatitude influences
- Dry-model experiment
- MJO-NAO two-way interactions

The extratropical influence on MJO is less well understood than the tropical influence on extratropics

# Examples of midlatitude influences

Ray and Zhang (2009)

Tropical channel model, two MJO events

The only factor found critical to the reproduction of the MJO initiation is time-varying lateral boundary conditions from the reanalysis. When such lateral boundary conditions are replaced by time-independent conditions, the model fails to reproduce the MJO initiation. These results support the idea that extratropical influences can be an efficient mechanism for MJO initiation.

Ray and Zhang (2010), importance of latitudinal momentum transport

# Examples of midlatitude influences

Hong et al. (2017),

extratropical forcing of 2015 MJO – El Nino event,  
southward penetration of north wind anomalies associated  
with extratropical disturbances in the extratropical western  
North Pacific

Nick Hall's next talk

# **Tropical-extratropical interactions in a dry GCM**

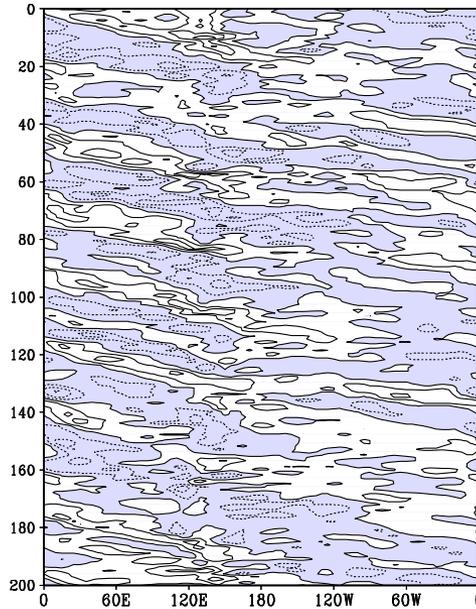
# Model and experiment

- Primitive equation AGCM (Hall 2000)
- T31, 10 levels
- Time-independent forcing to maintain the winter climate  
→ all variabilities come from internal dynamics
- No moisture equation, no interactive convection
- 3660 days of perpetual winter integration

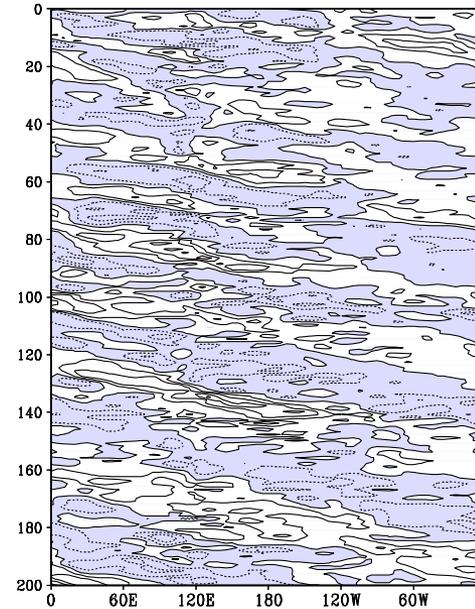
Zonal propagation  
10S-10N

Unfiltered data

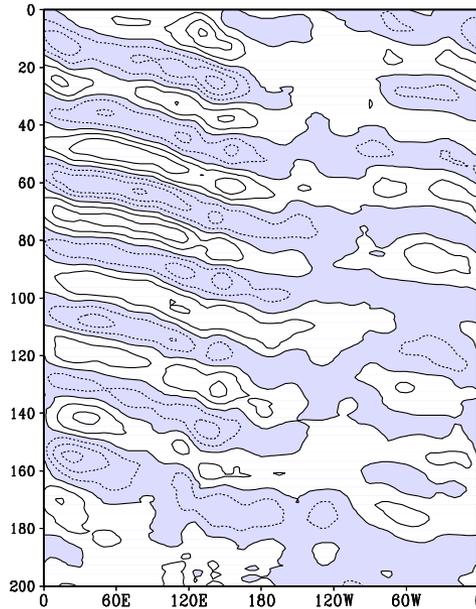
a) VP250 from day500



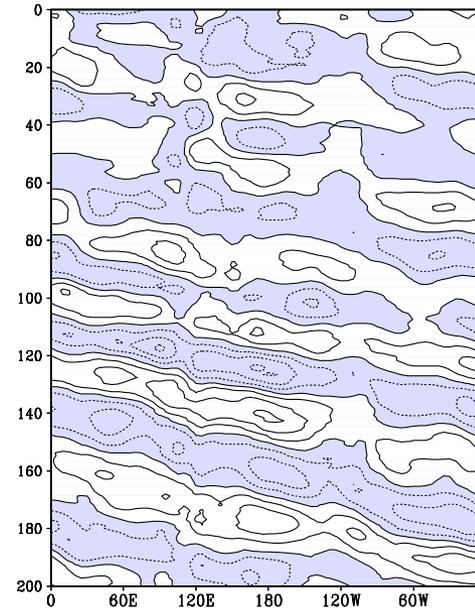
b) VP250 from day2800



c) LF VP250 from day500



d) LF VP250 from day2800

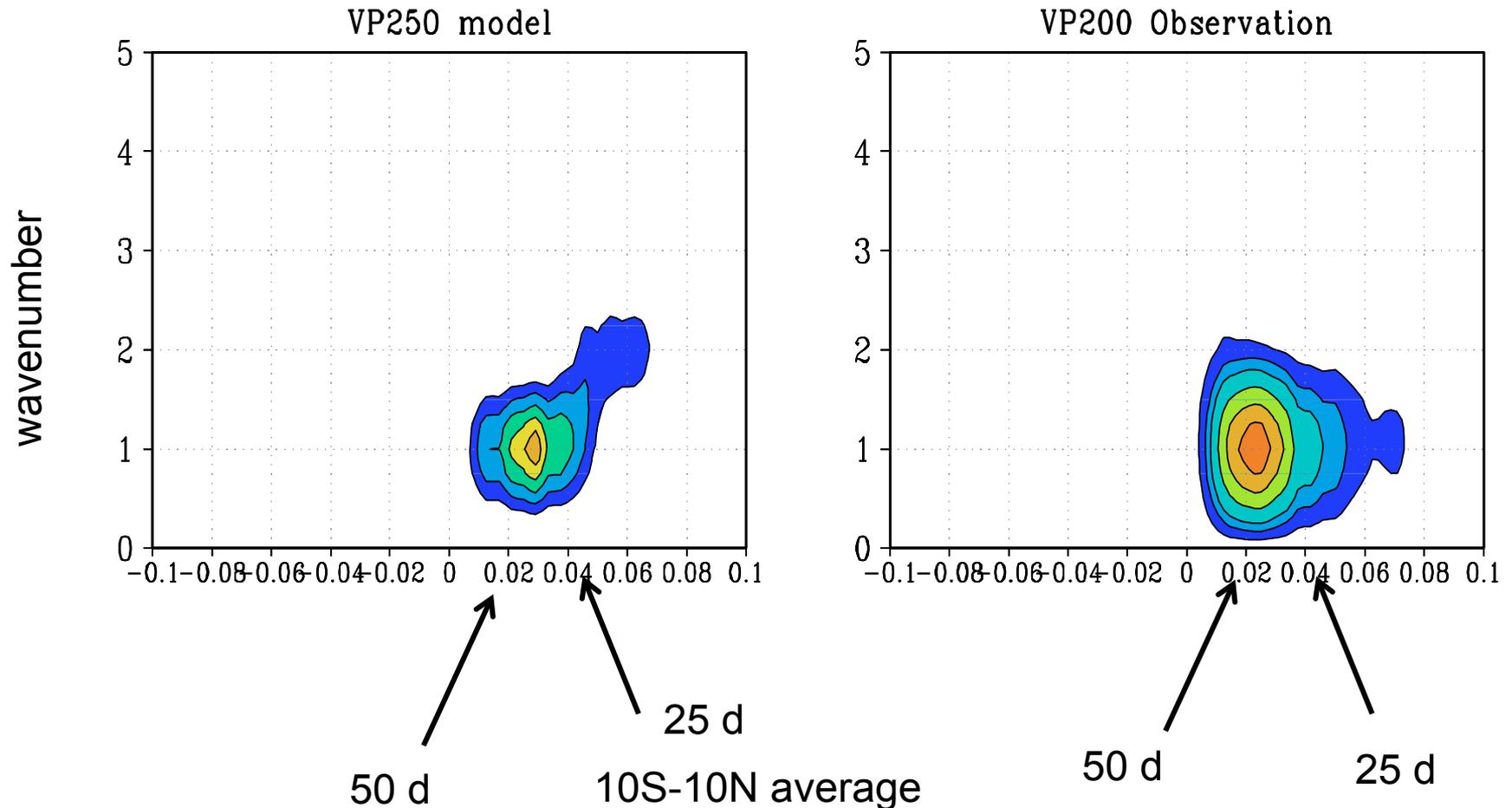


20-100 day band-pass

Stronger in eastern Hemisphere

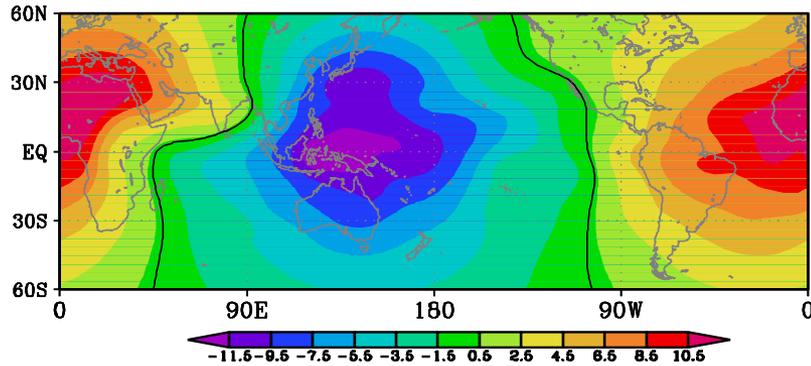
# Wavenumber-frequency spectra

## Equatorial velocity potential

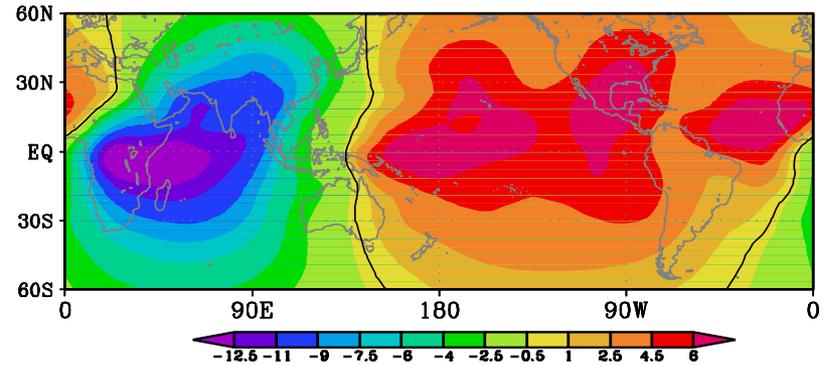


# EOF analysis of 20-100 day band-passed 250 hPa velocity potential

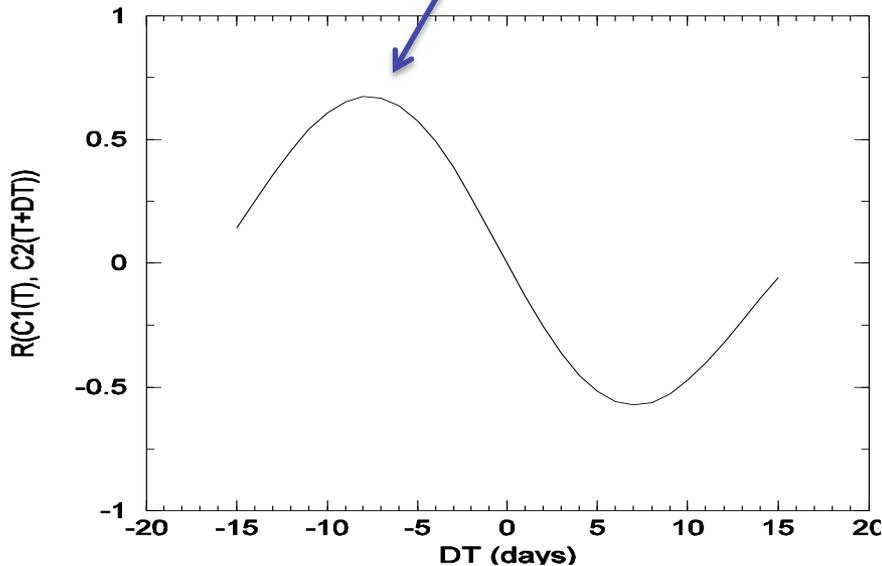
a) EOF1 34%



b) EOF2 30%

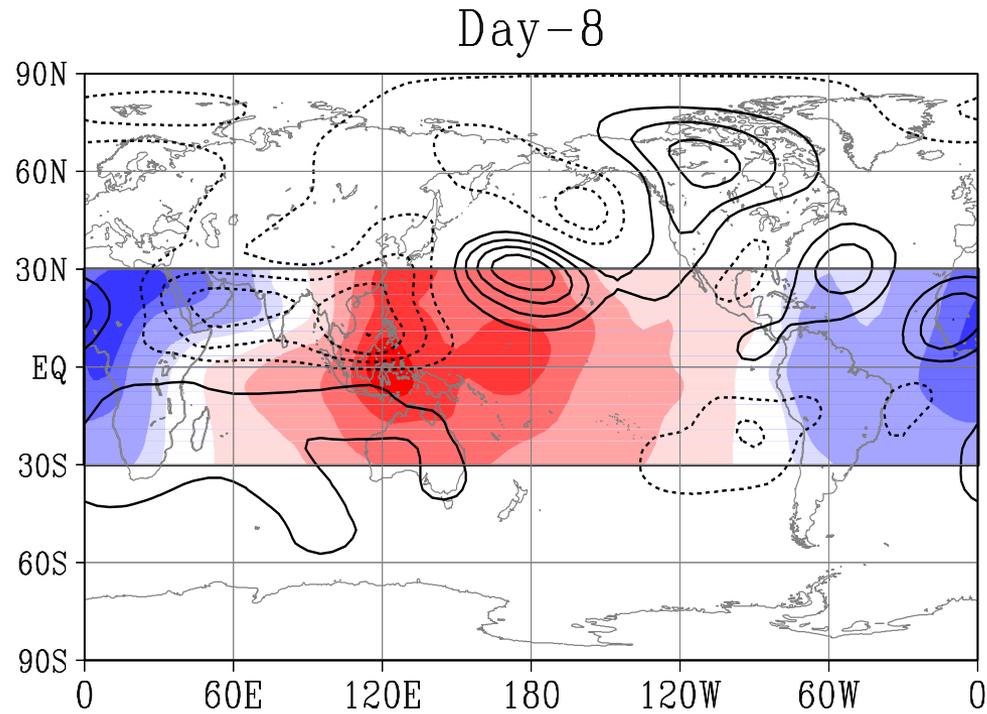


PC2 lead



TIV index:

$$I(t) = \frac{PC_2(t) + PC_1(t + 8)}{2}$$

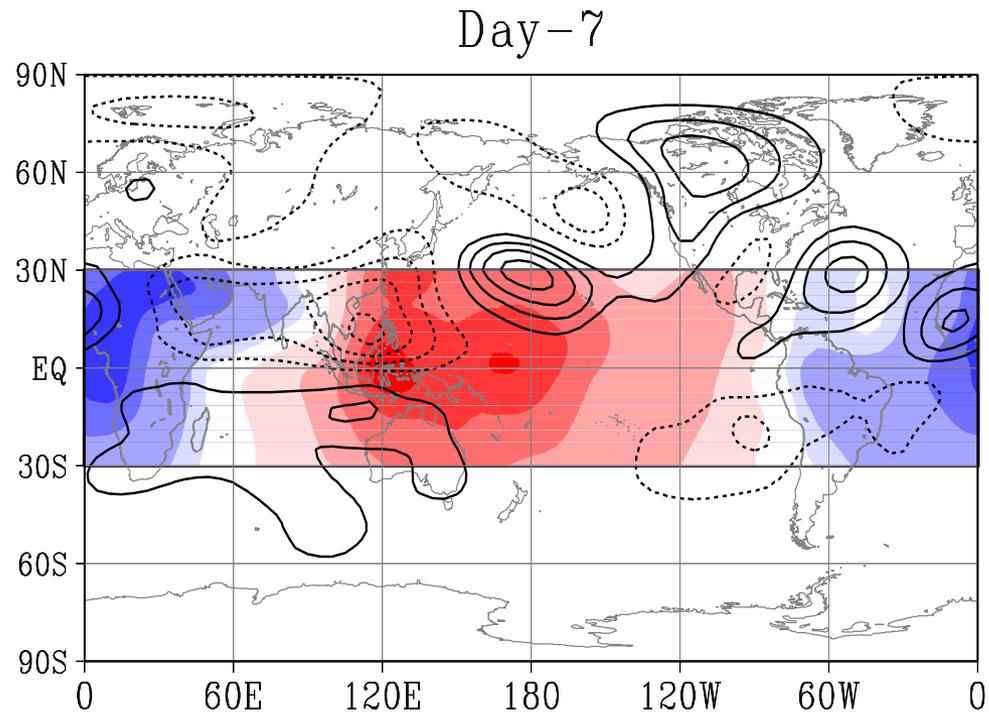


## Regression to TIV index

Color: 250mb velocity potential

Contour: 250mb streamfunction anomaly

**TIV index:** In phase with PC2

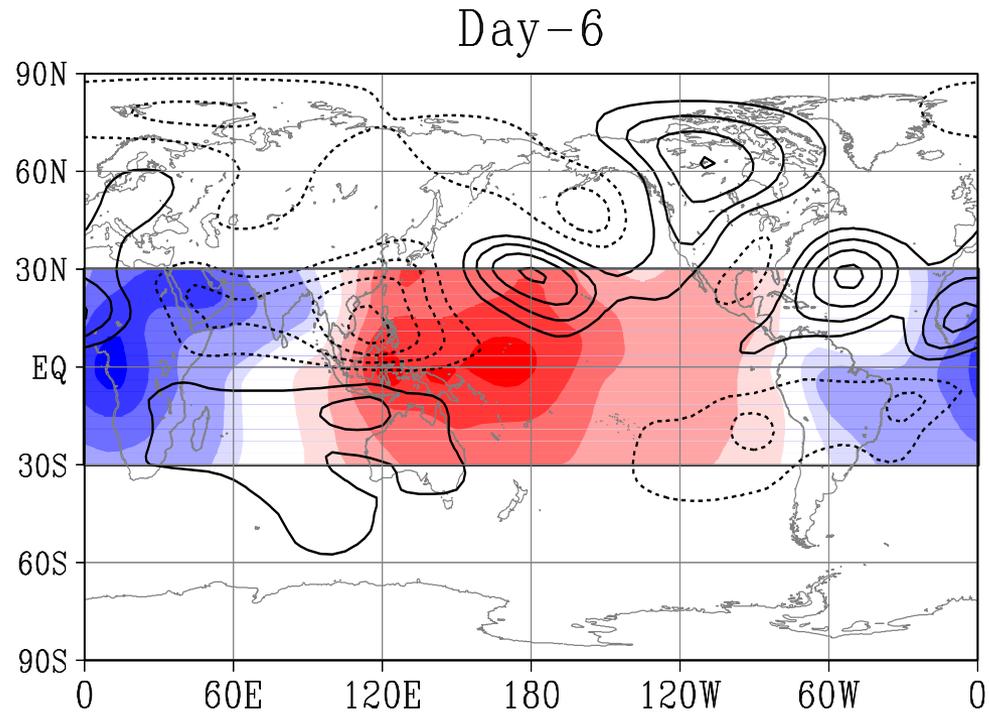


## Regression to TIV index

Color: 250mb velocity potential

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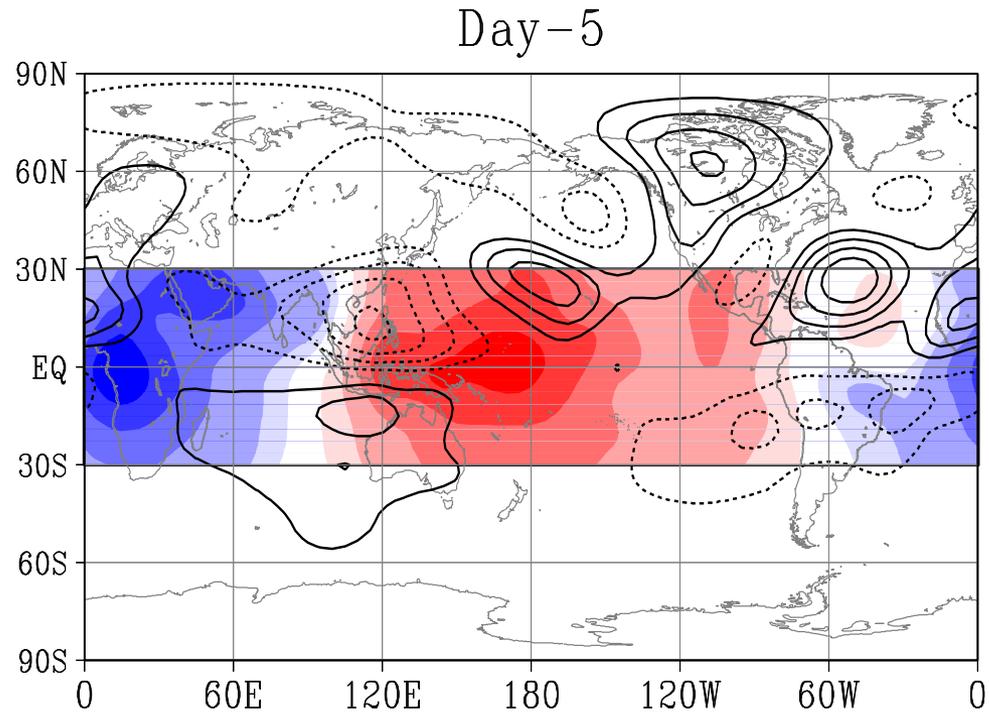


## Regression to TIV index

Color: 250mb velocity potential

Contour: 250mb streamfunction anomaly

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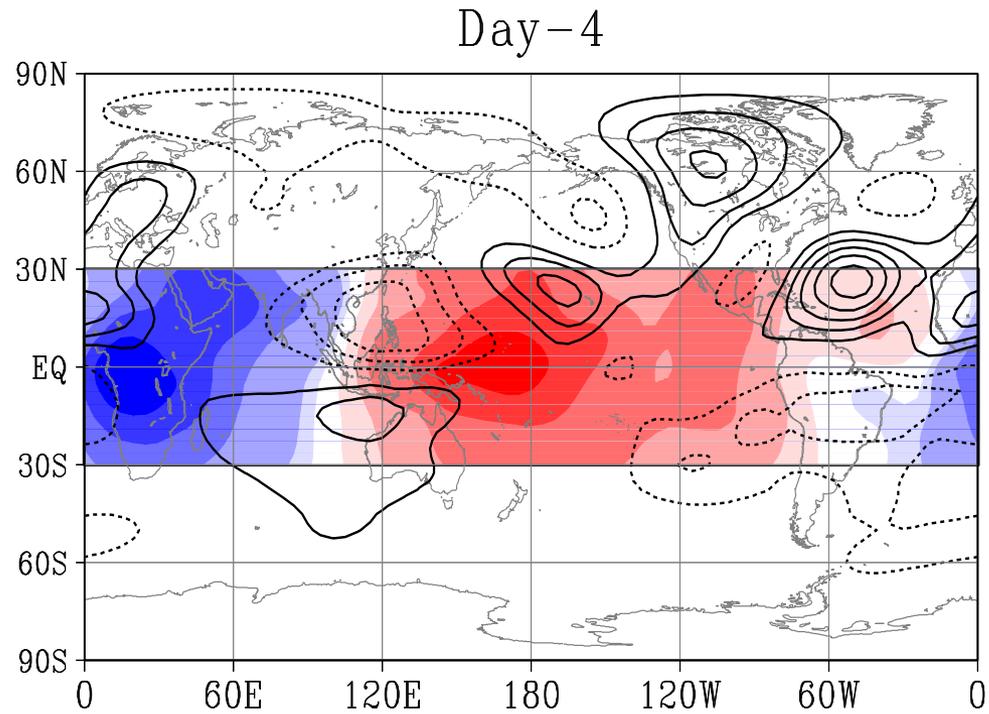


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TIV index: In phase with PC2

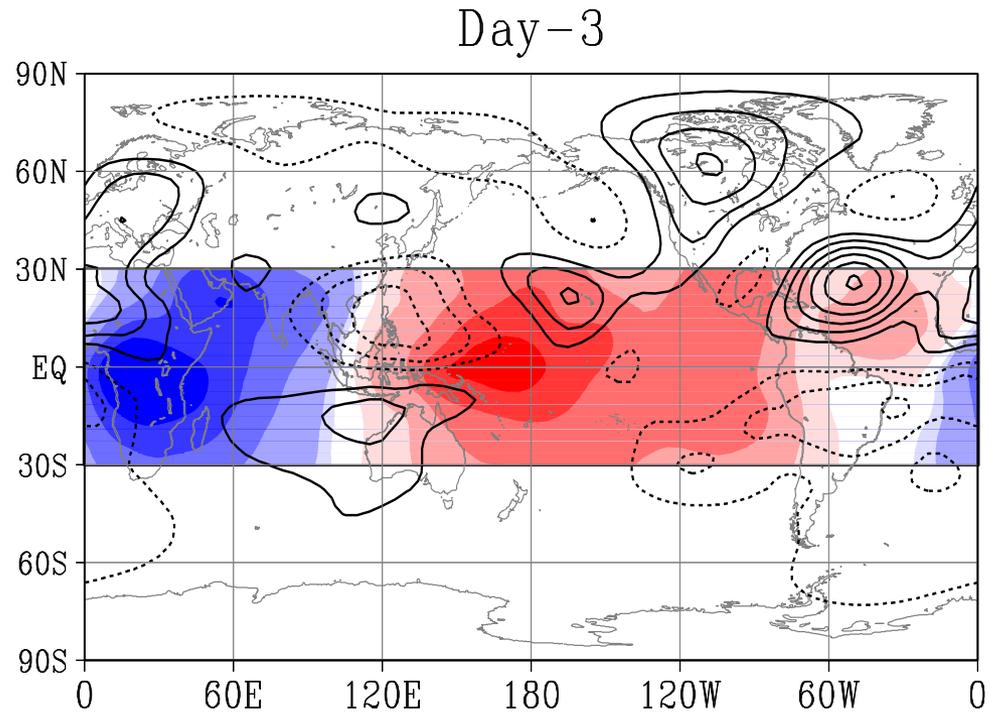


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**TIV index:** In phase with PC2

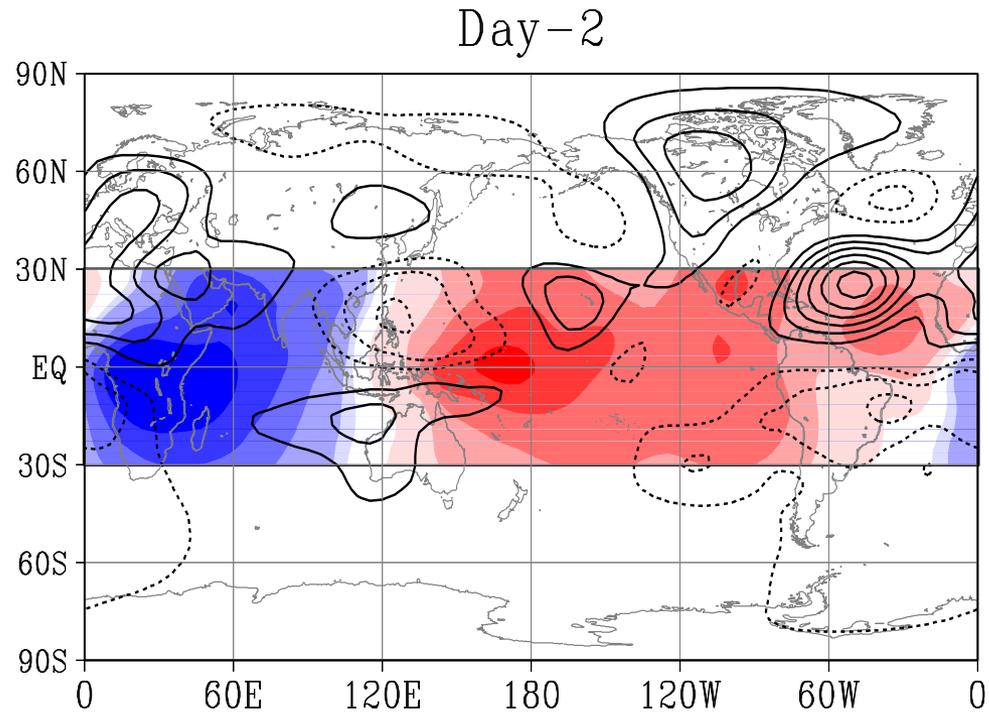


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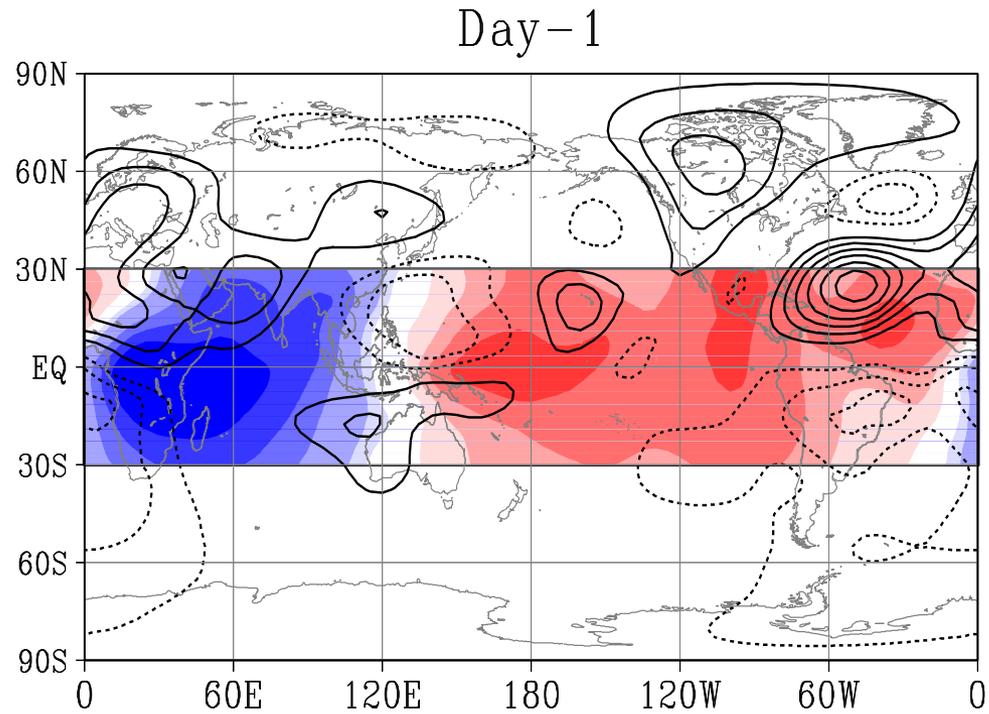


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Color: 250mb velocity potential

Contour: 250mb streamfunction anomaly

TIV index: In phase with PC2

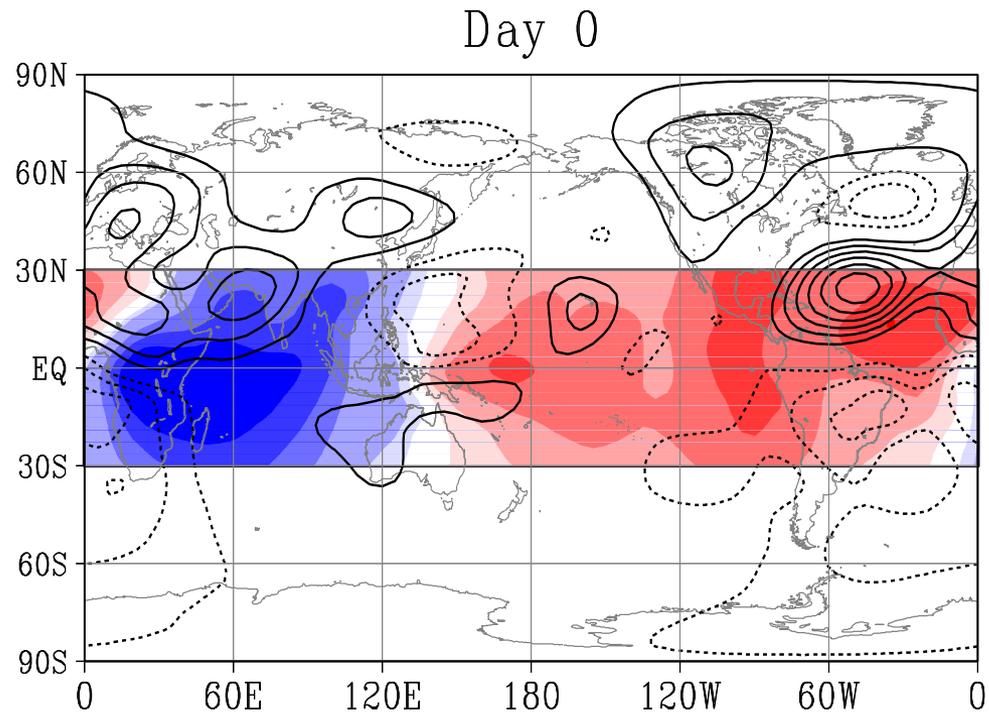


## Regression to TIV index

Color: 250mb velocity potential

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**TIV index:** In phase with PC2

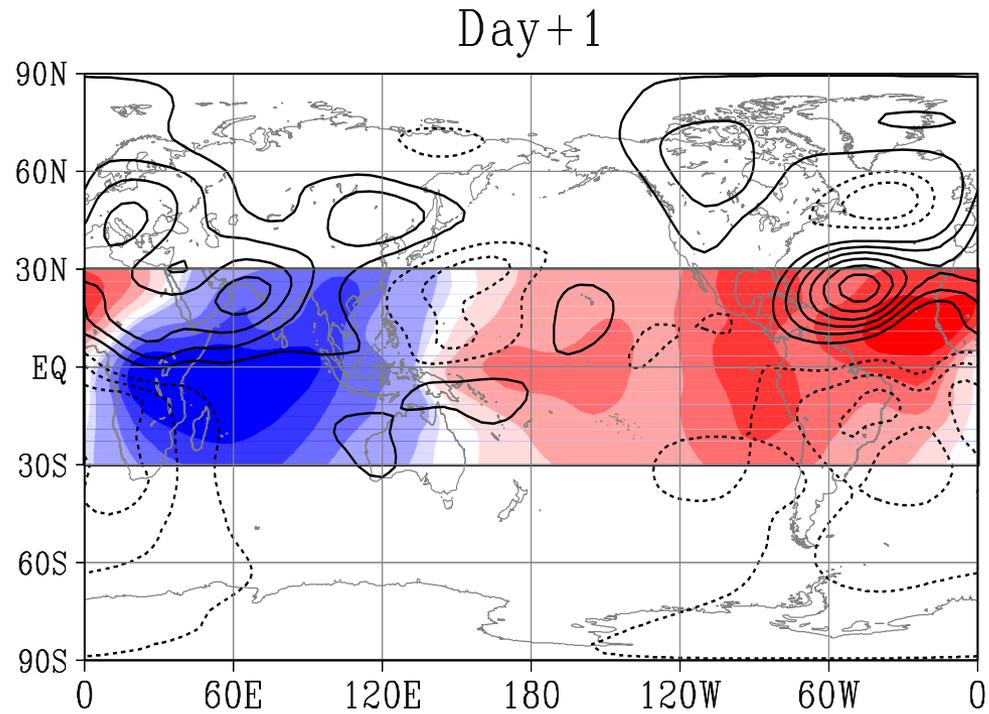


## Regression to TIV index

Color: 250mb velocity potential

Contour: 250mb streamfunction anomaly

TIV index: In phase with PC2

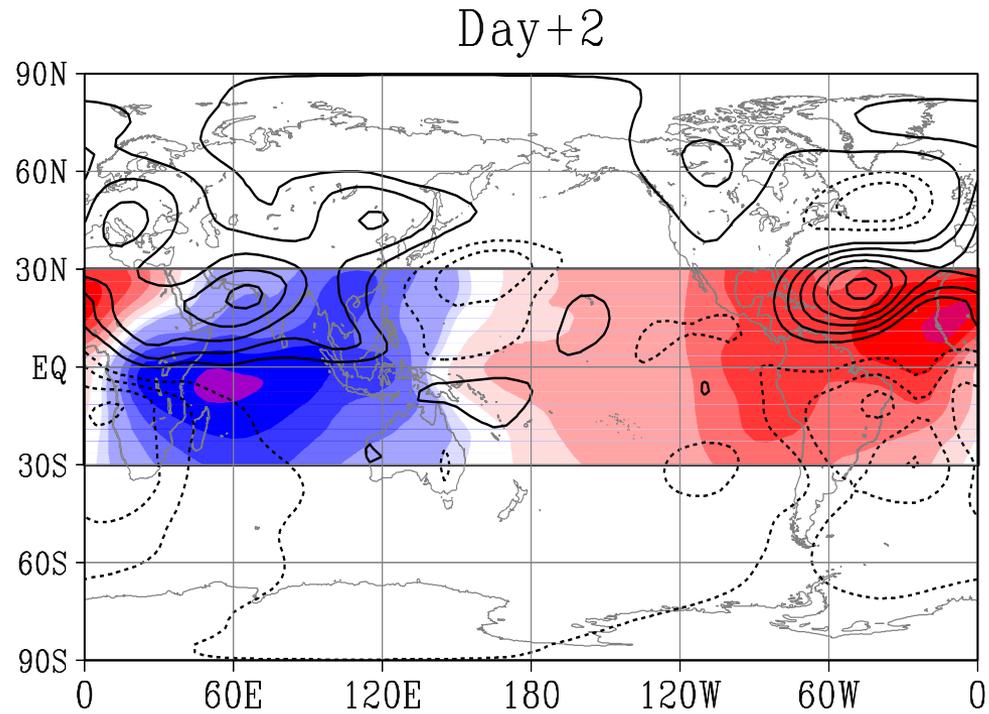


## Regression to TIV index

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Contour: 250mb streamfunction anomaly

**TIV index:** In phase with PC2

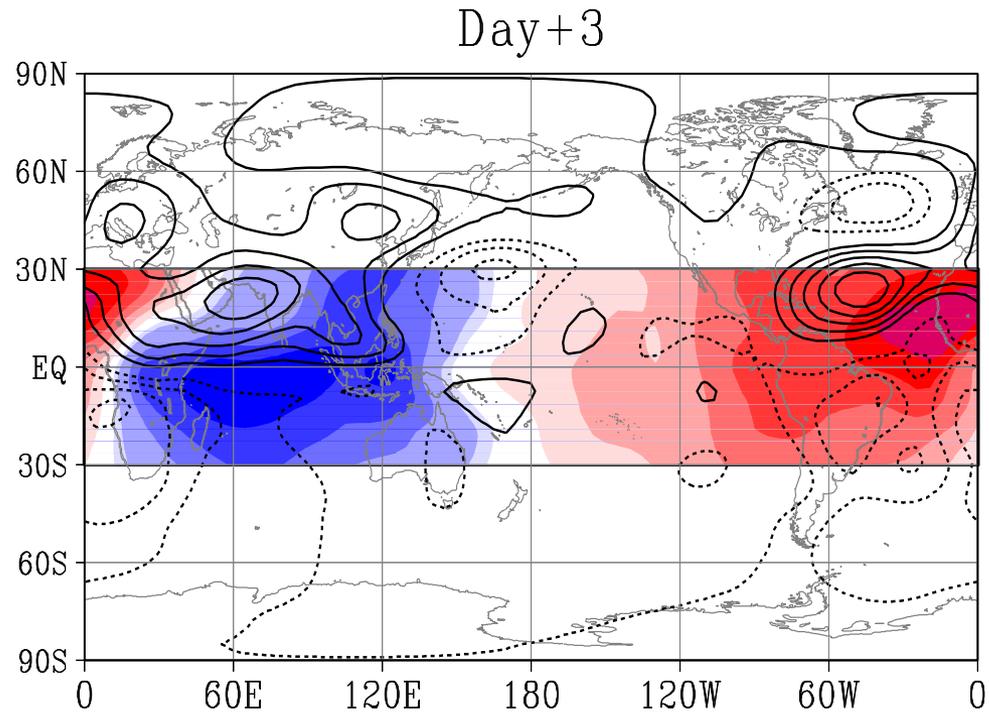


## Regression to TIV index

Color: 250mb velocity potential

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TIV index: In phase with PC2

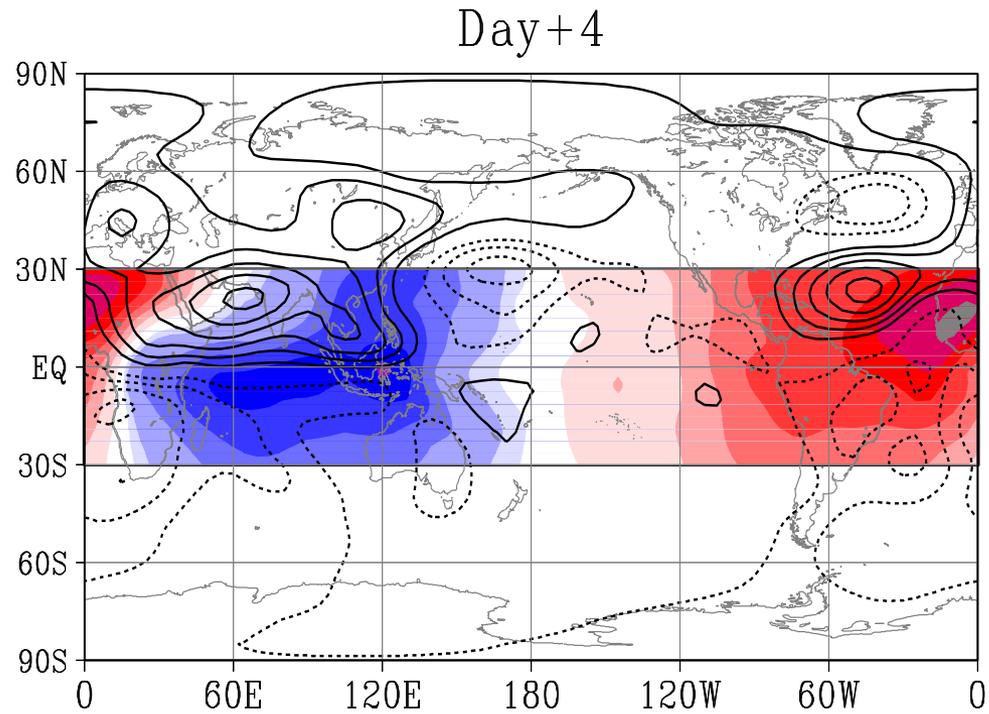


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Color: 250mb velocity potential

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TIV index: In phase with PC2

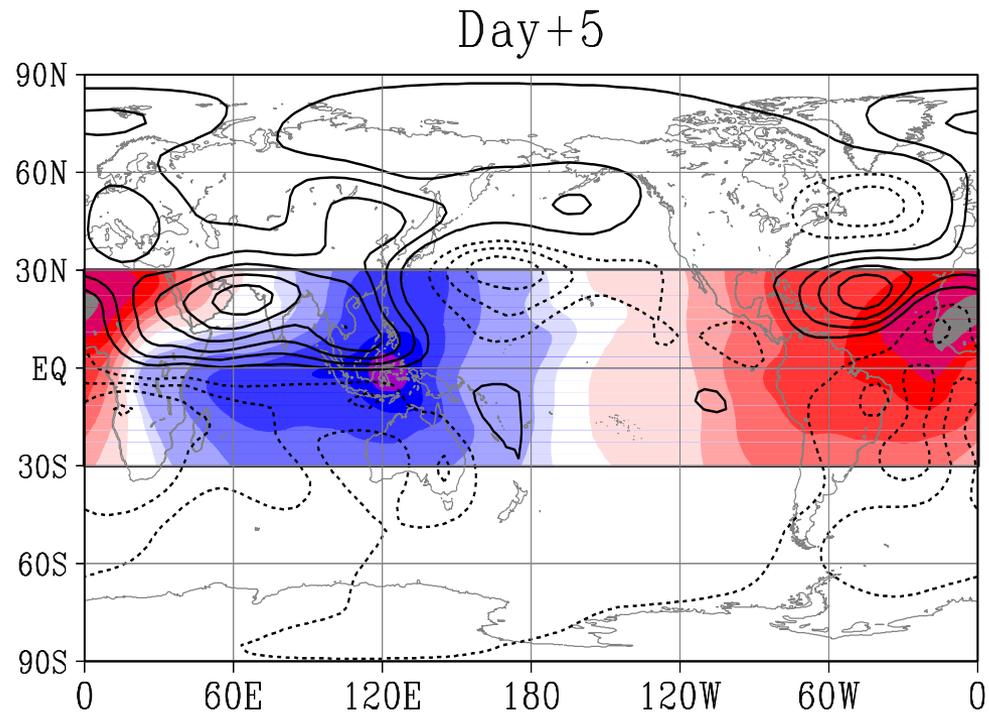


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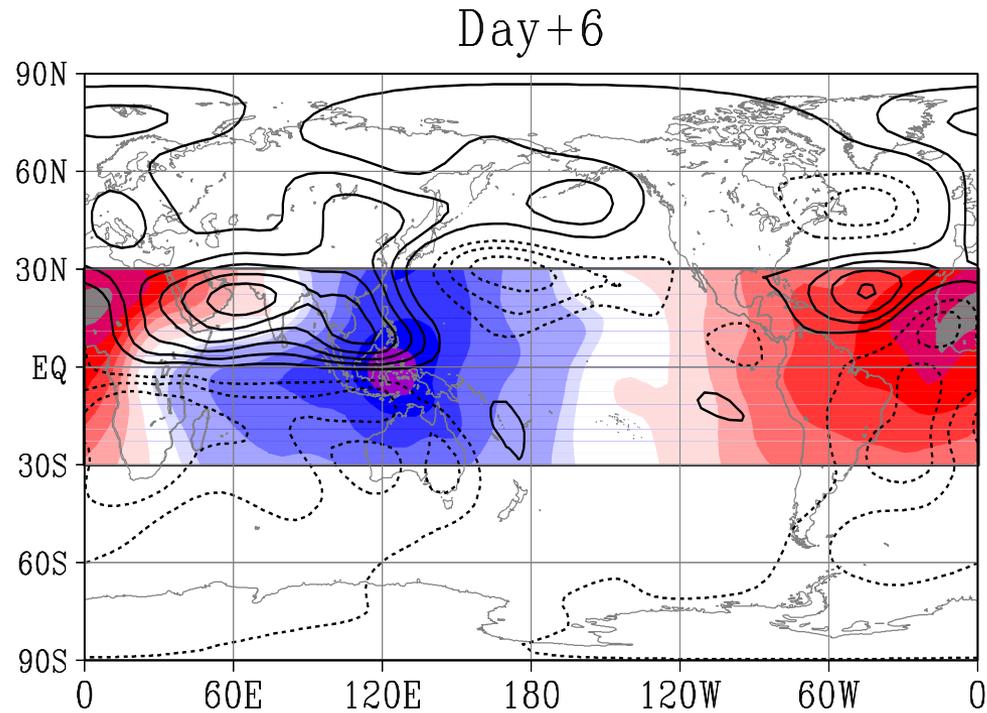


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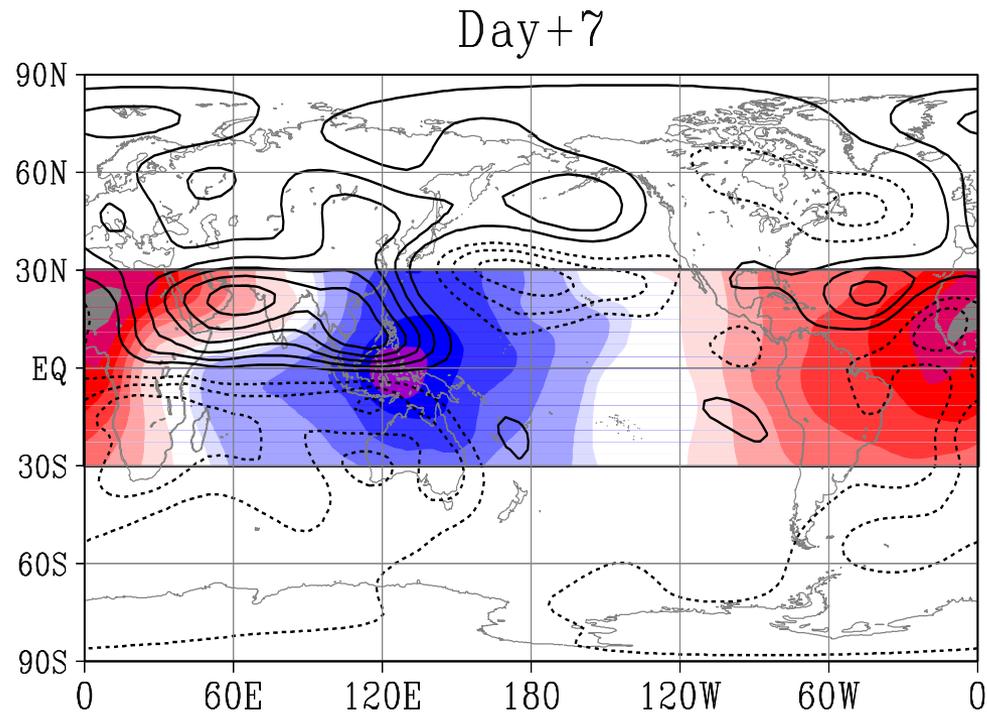


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Color: 250mb velocity potential

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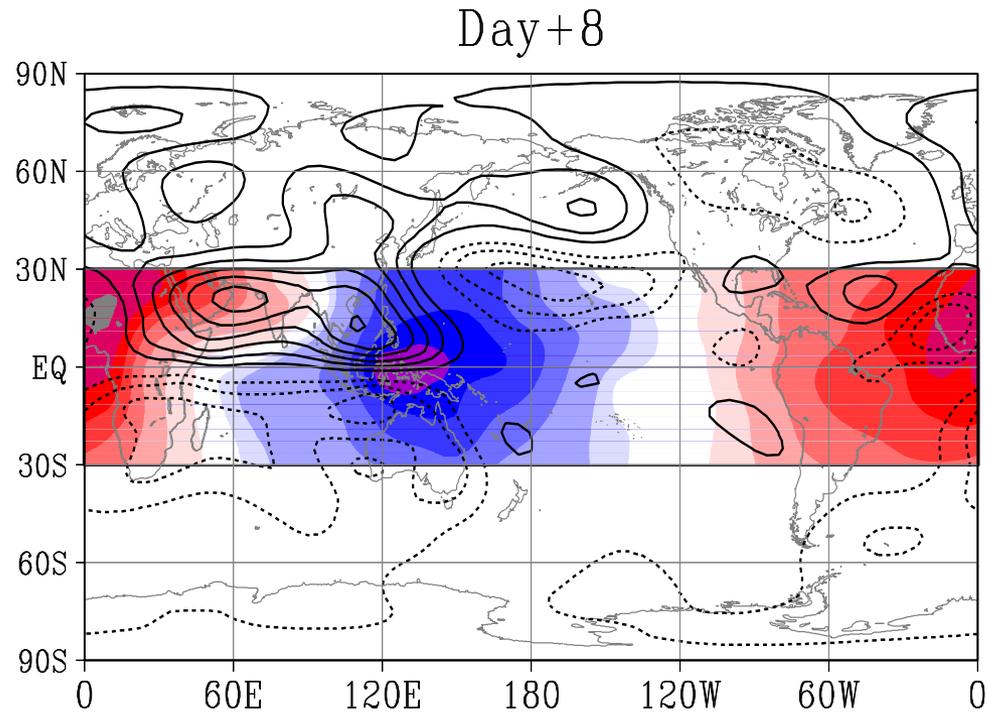


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Color: 250mb velocity potential

Contour: 250mb streamfunction anomaly

TIV index: In phase with PC2

# Waveflux - Theory

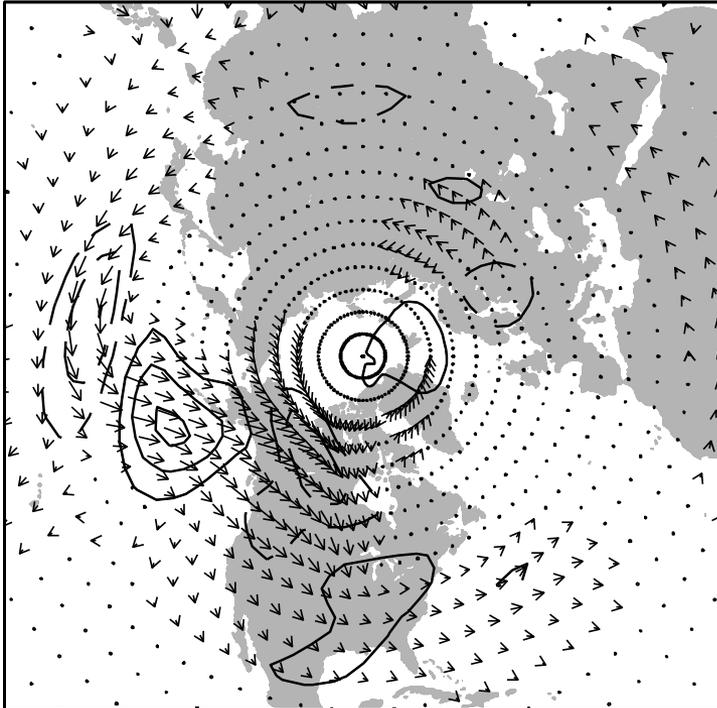
$$\mathbf{W} = \frac{1}{2|\mathbf{U}|} \begin{bmatrix} U(\psi_x^2 - \psi\psi_{xx}) + V(\psi_x\psi_y - \psi\psi_{xy}) \\ U(\psi_x\psi_y - \psi\psi_{xy}) + V(\psi_y^2 - \psi\psi_{yy}) \end{bmatrix}$$

Takaya and Nakamura 2001 GRL

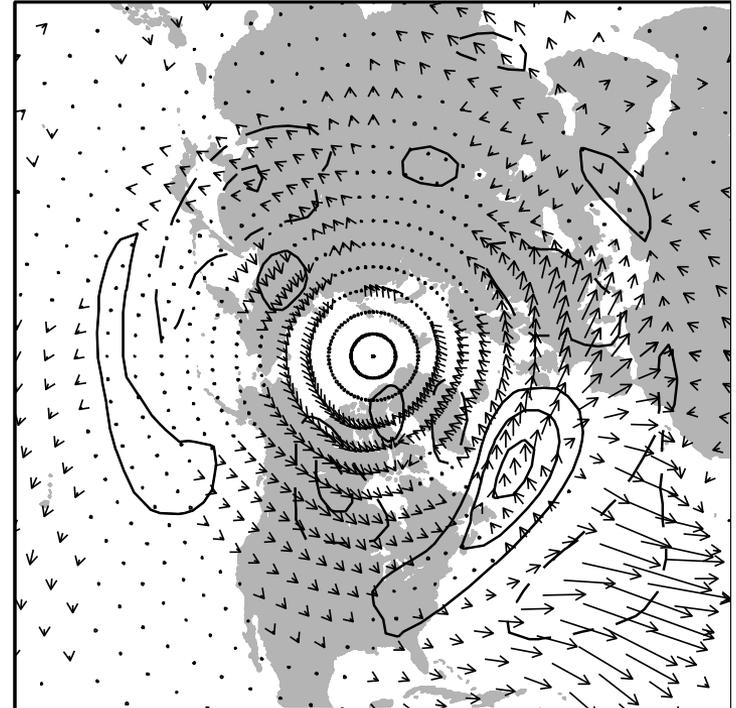
# ISO in a dry model

250 hPa PV' and wave activity flux

a) Day -8



b) Day 0



Linked to tropical eastward propagation in the eastern Hemisphere → Global propagation of low-frequency wave activity

# Summary

- TIV generated in a dry GCM
- Tropical-extratropical interactions are crucial in generating the model TIV
- Extratropical influence on tropical waves

Remaining questions:

- Contribution from moisture and convection
- Mechanism: how do extratropical large-scale disturbances, that are equivalent barotropic, propagate into the tropics to generate tropical waves that are baroclinic?

## **MJO-NAO two-way interactions**

# Data

NAO index: pentad average

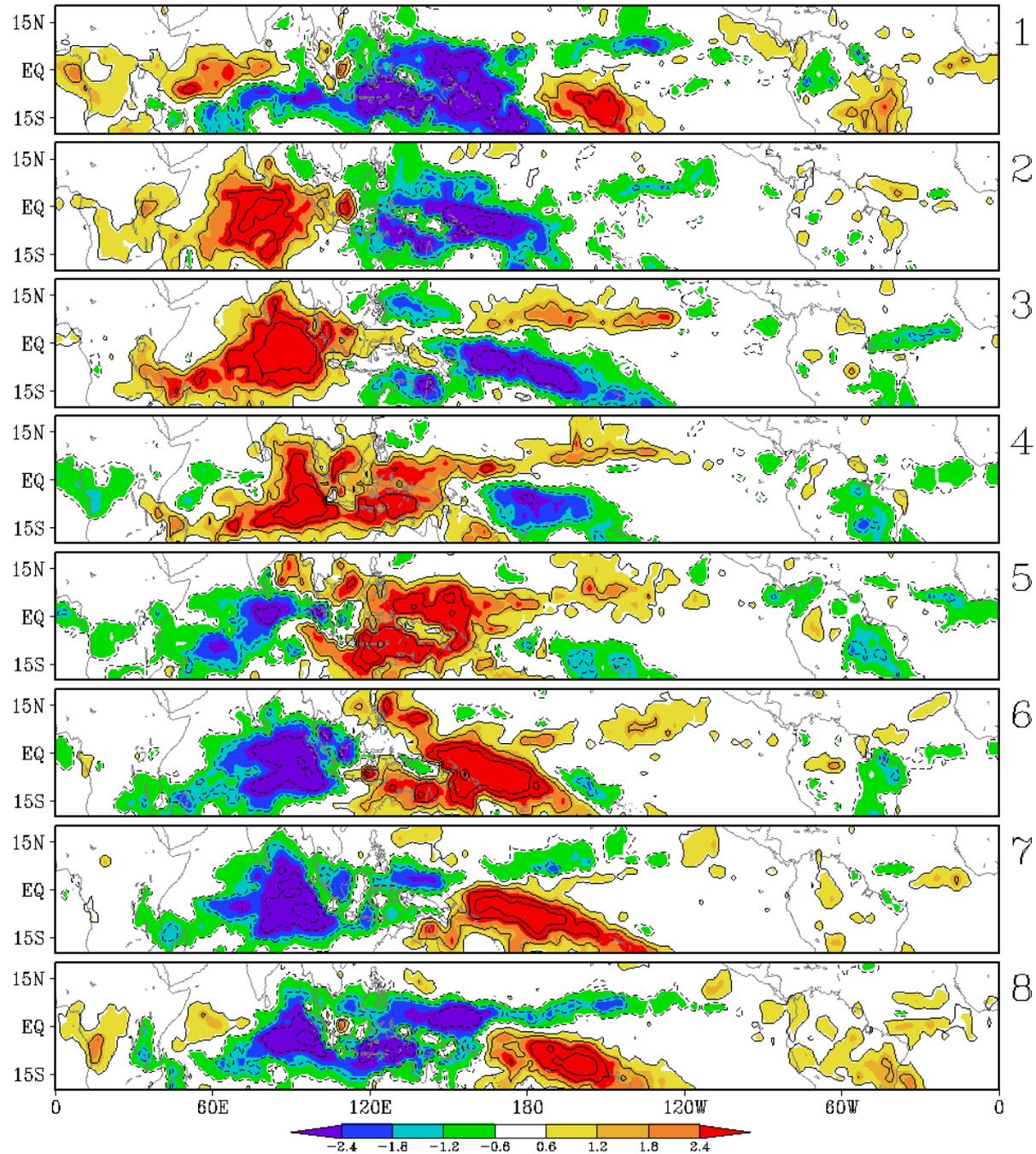
MJO RMMs: pentad average

Period: 1979-2003

Extended winter, November to April (36 pentads each winter)

Composites of tropical  
Precipitation rate for 8  
MJO phases,  
according to Wheeler  
and Hendon index.

Xie and Arkin pentad  
data, 1979-2003



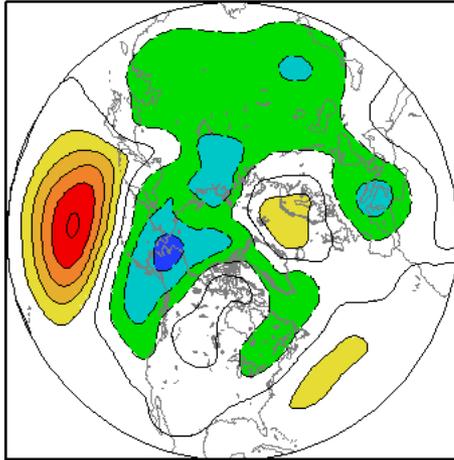
# Lagged probability of the NAO index

Positive: upper tercile; Negative: low tercile

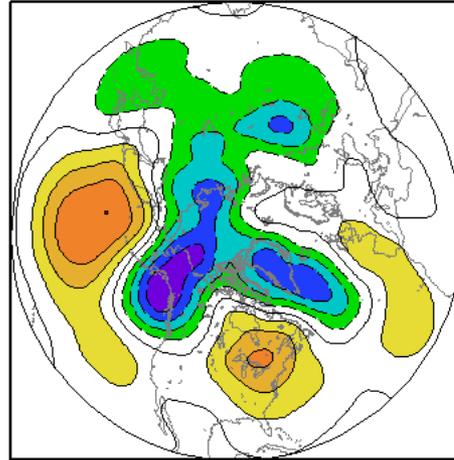
Phase	1	2	3	4	5	6	7	8
Lag -5		-35%	-40%			+49%	+49%	
Lag -4						+52%	+46%	
Lag -3		-40%					+46%	
Lag -2						+50%		
Lag -1								
Lag 0				+45%				-42%
Lag +1			+47%	+45%				-46%
Lag +2		+47%	+50%	+42%		-41%	-41%	-42%
Lag +3		+48%				-41%	-48%	
Lag +4						-39%	-48%	
Lag +5				-41%				

# Tropical influence

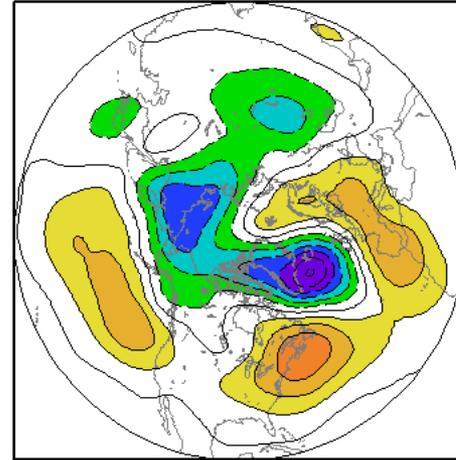
a) PHASE 3 lag=0



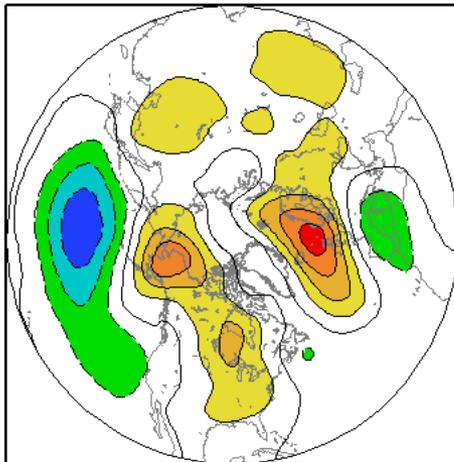
b) PHASE 3 lag=1



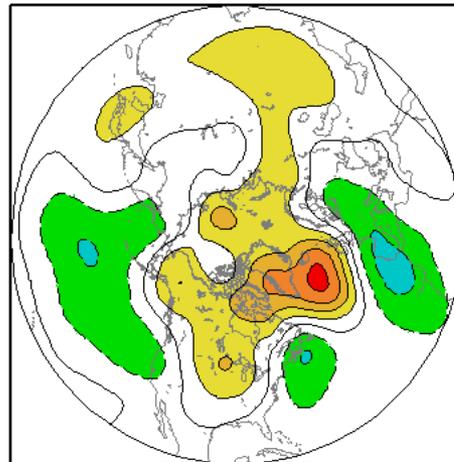
c) PHASE 3 lag=2



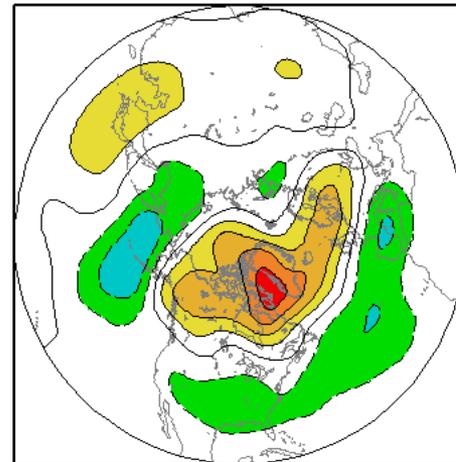
d) PHASE 7 lag=0



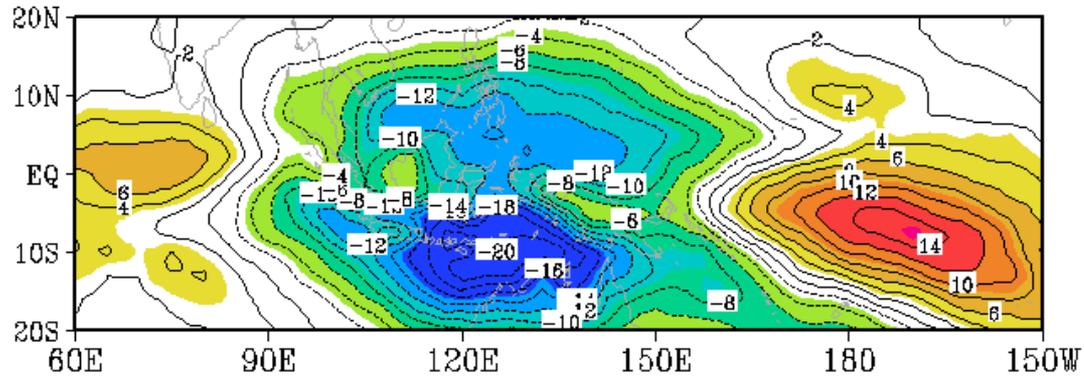
e) PHASE 7 lag=1



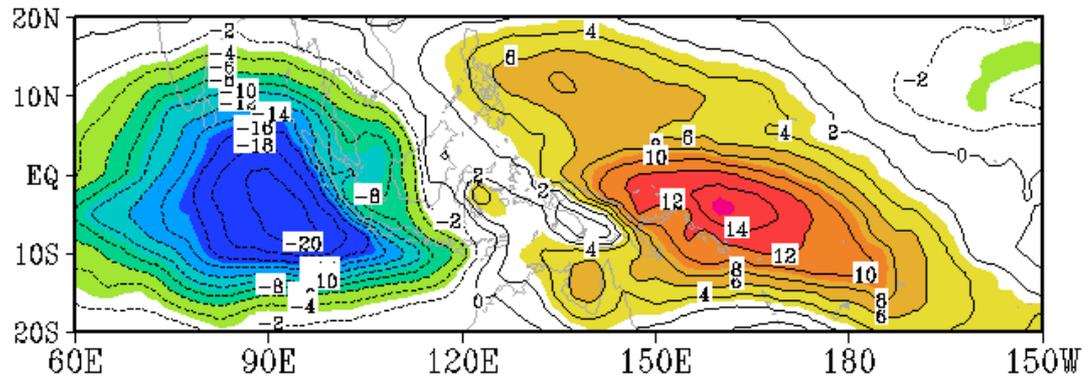
f) PHASE 7 lag=2



a) OLR EOF1 11%

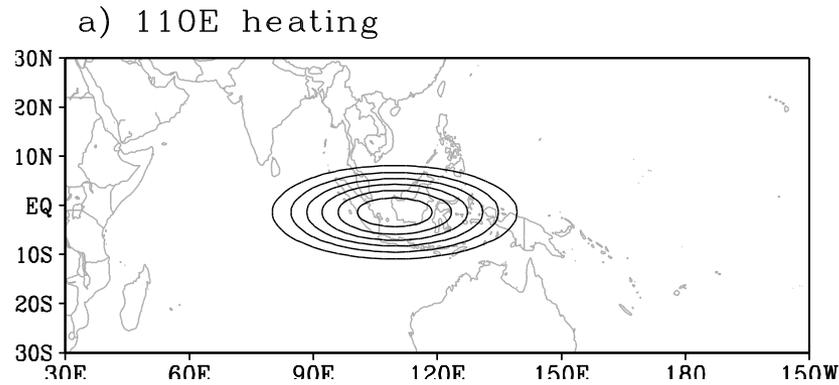


b) OLR EOF2 10%

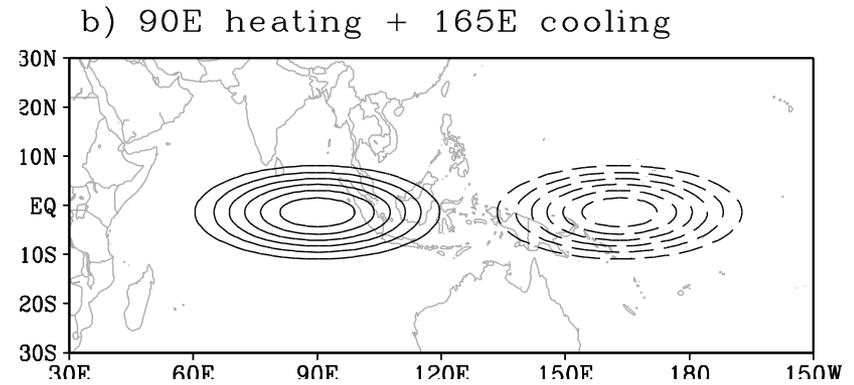


Correlation when PC2 leads PC1 by 2 pentads: 0.66

# Thermal forcing



Exp1 forcing



Exp2 forcing

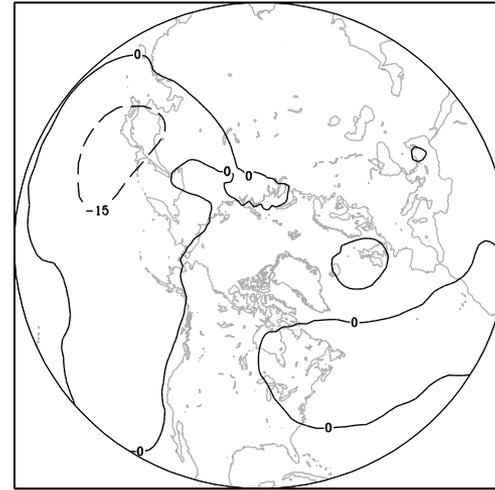
# Z500 response

Exp1

a) Exp1: days6-10

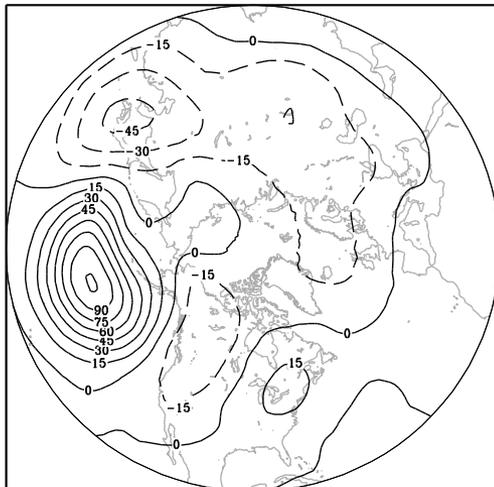


b) Exp1: days11-15

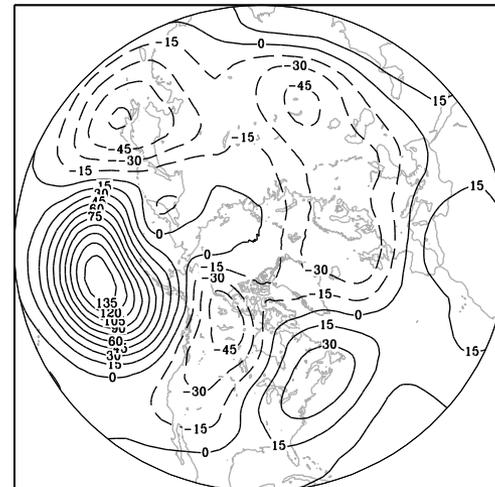


Exp2

c) Exp2: days6-10



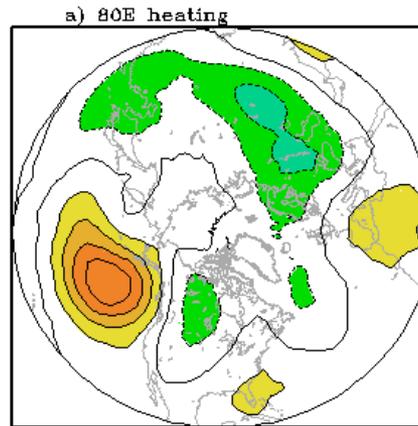
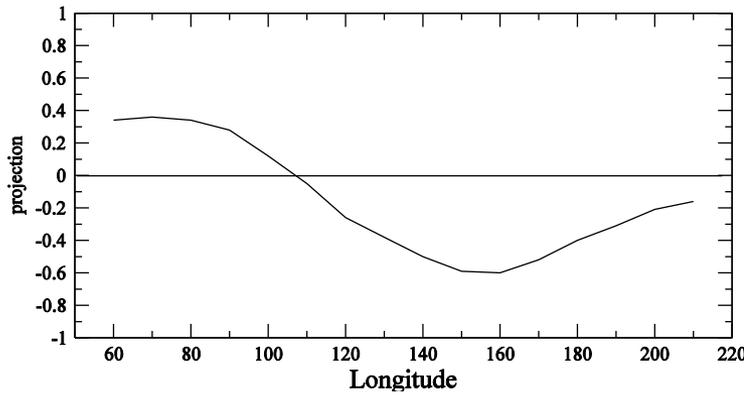
d) Exp2: days11-15



## Why the response to a dipole heating is the strongest ?

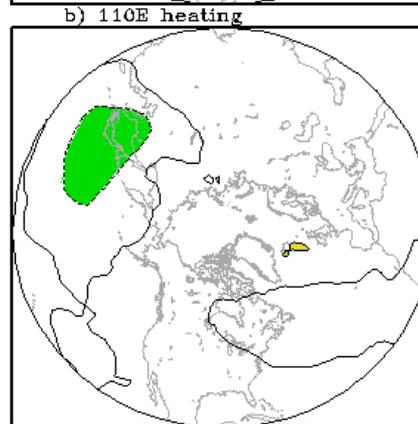
- Linear integration, winter basic state
- with a single center heating source
- Heating at different longitudes along the equator from 60E to 150W at a 10 degree interval, 16 experiments
- Z500 response at day 10

# Day 10 Z500 linear response

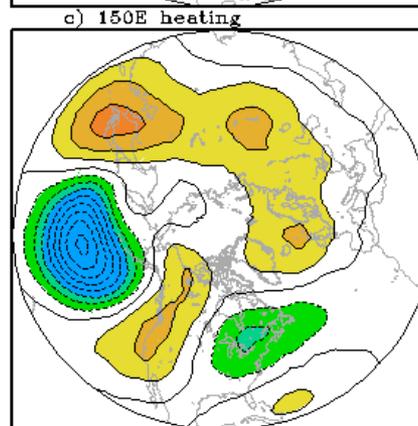


80E

Similar pattern for heating 60-100E



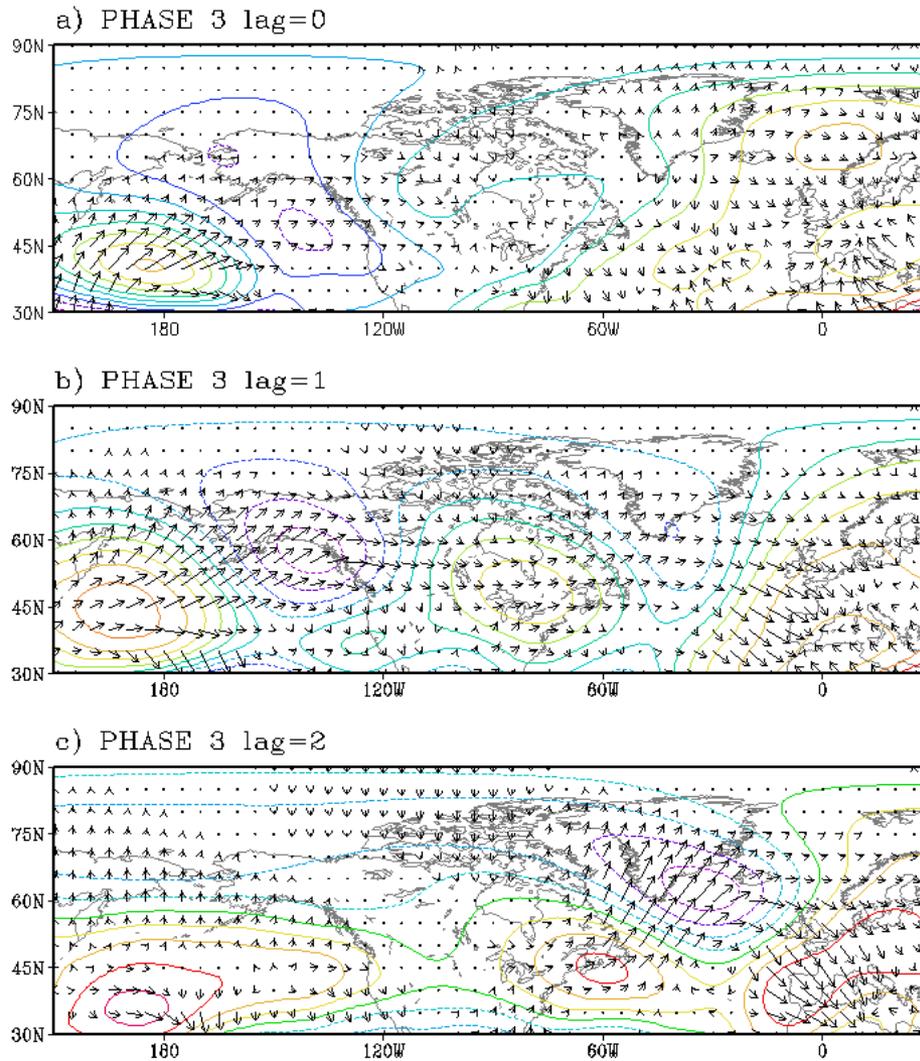
110E



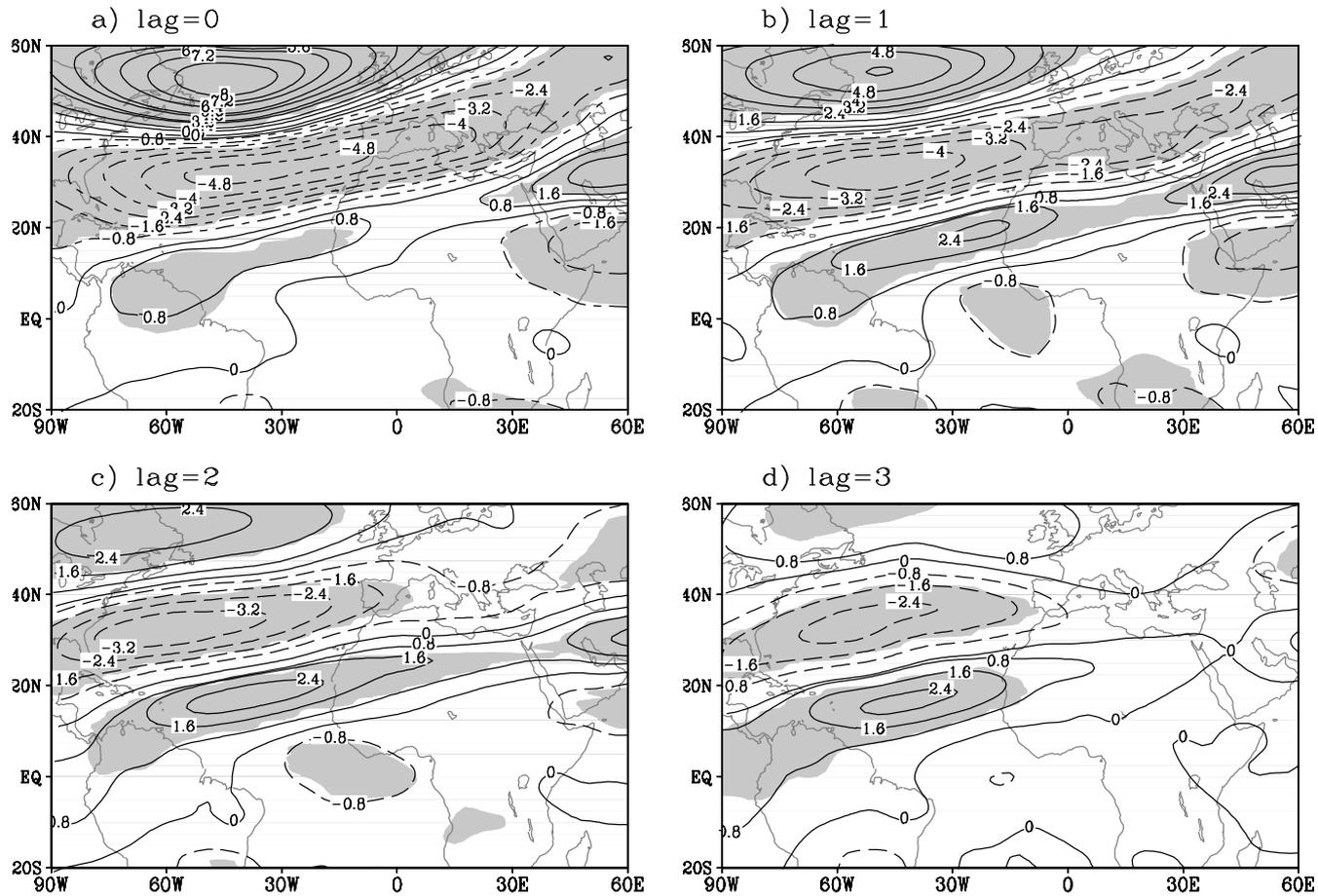
150E

Similar pattern for heating 120-150W

# Wave activity flux and 200mb streamfunction anomaly

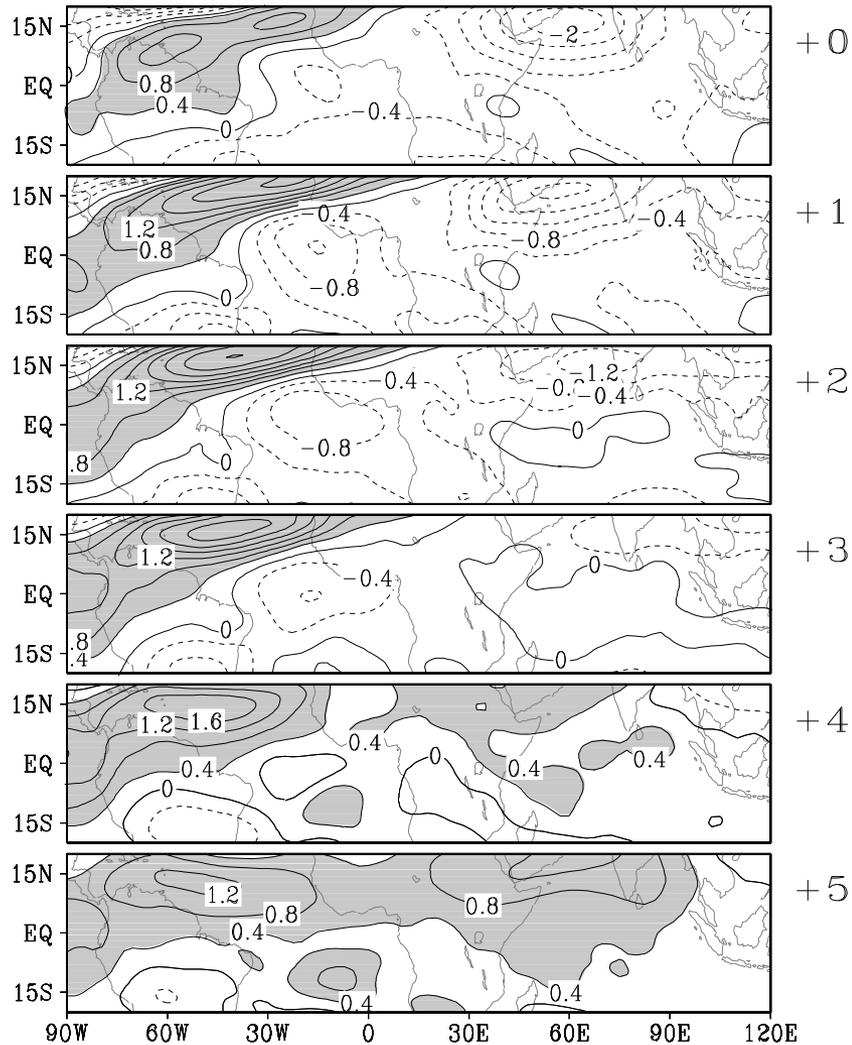


# Extratropical influence



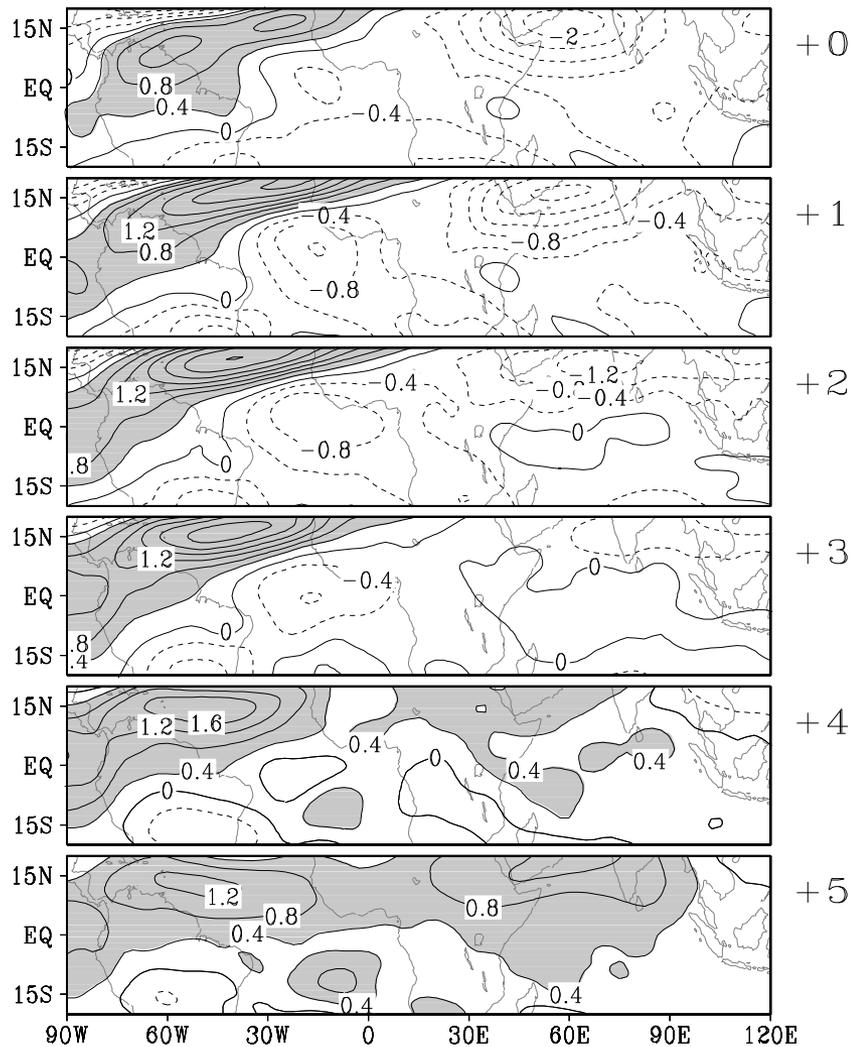
Lagged regression of 200mb U to NAO index

# Extratropical influence

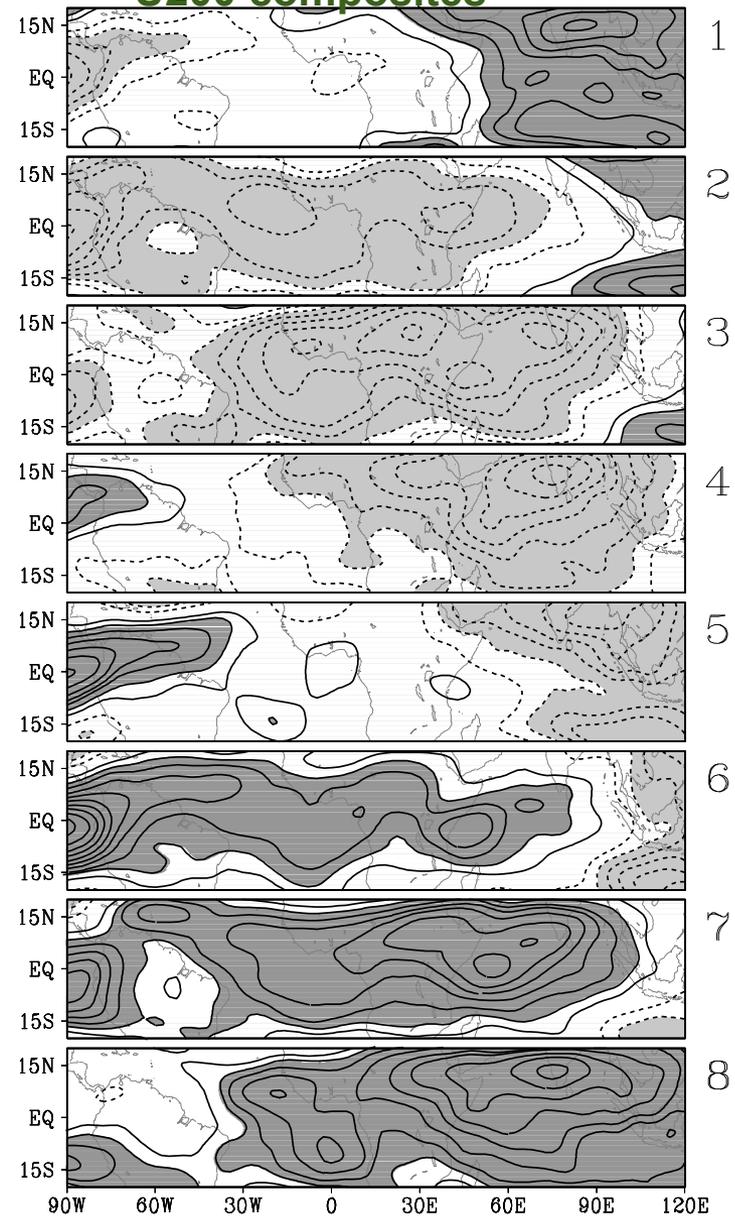


Lagged regression of 200mb U to NAO index

# Extratropical influence

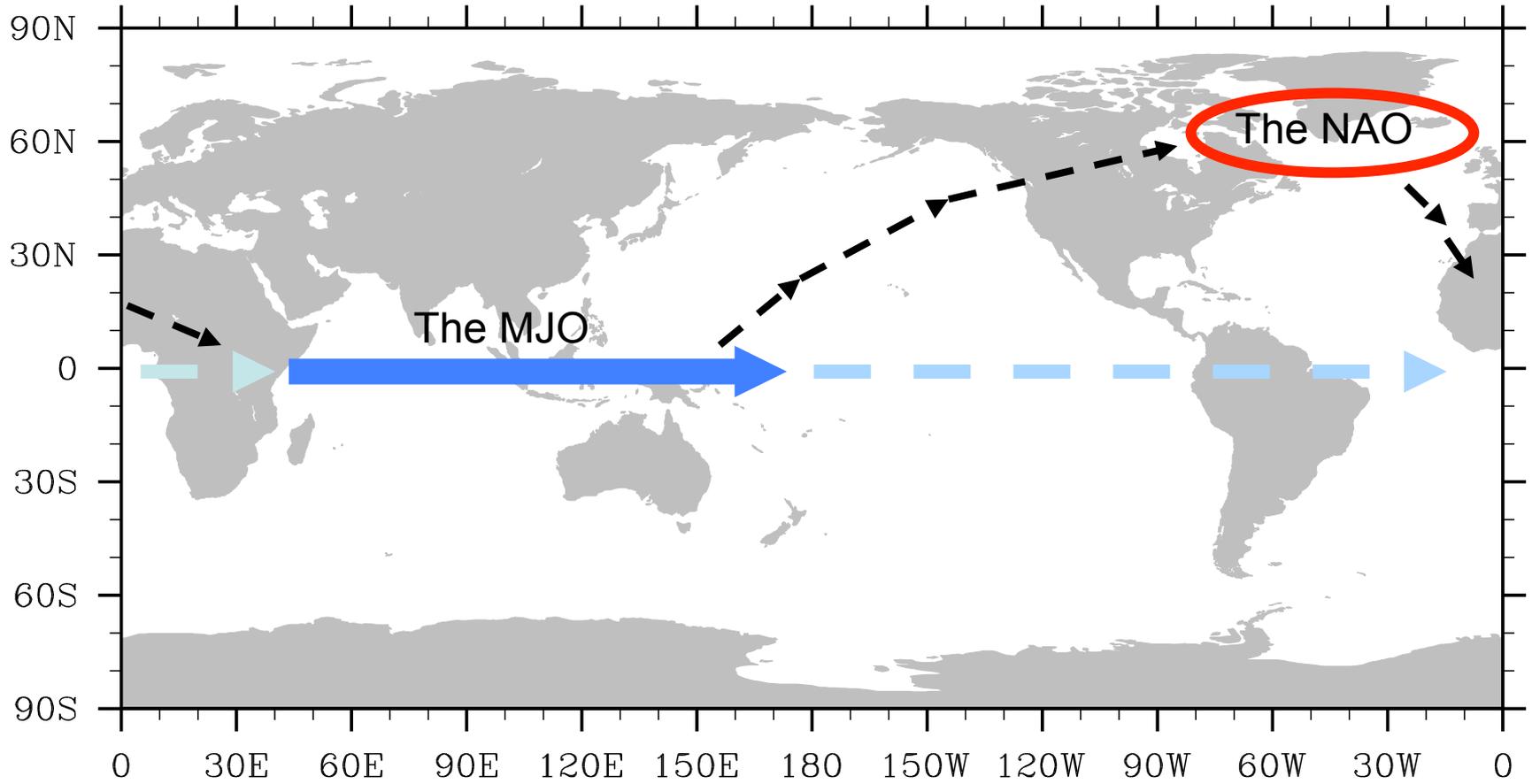


## U200 composites



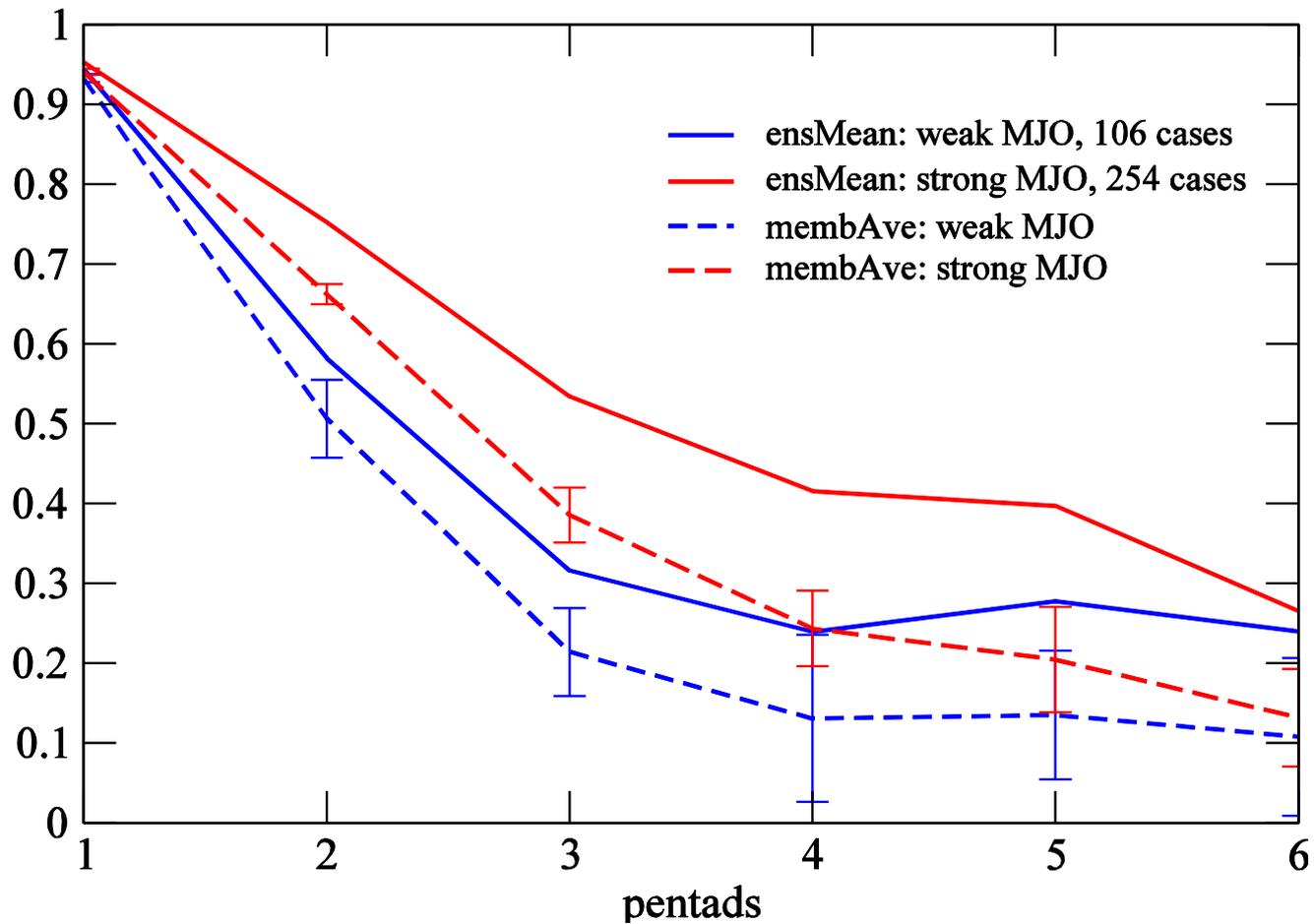
Lagged regression of 200mb U to NAO index

# Two-way MJO – NAO interaction

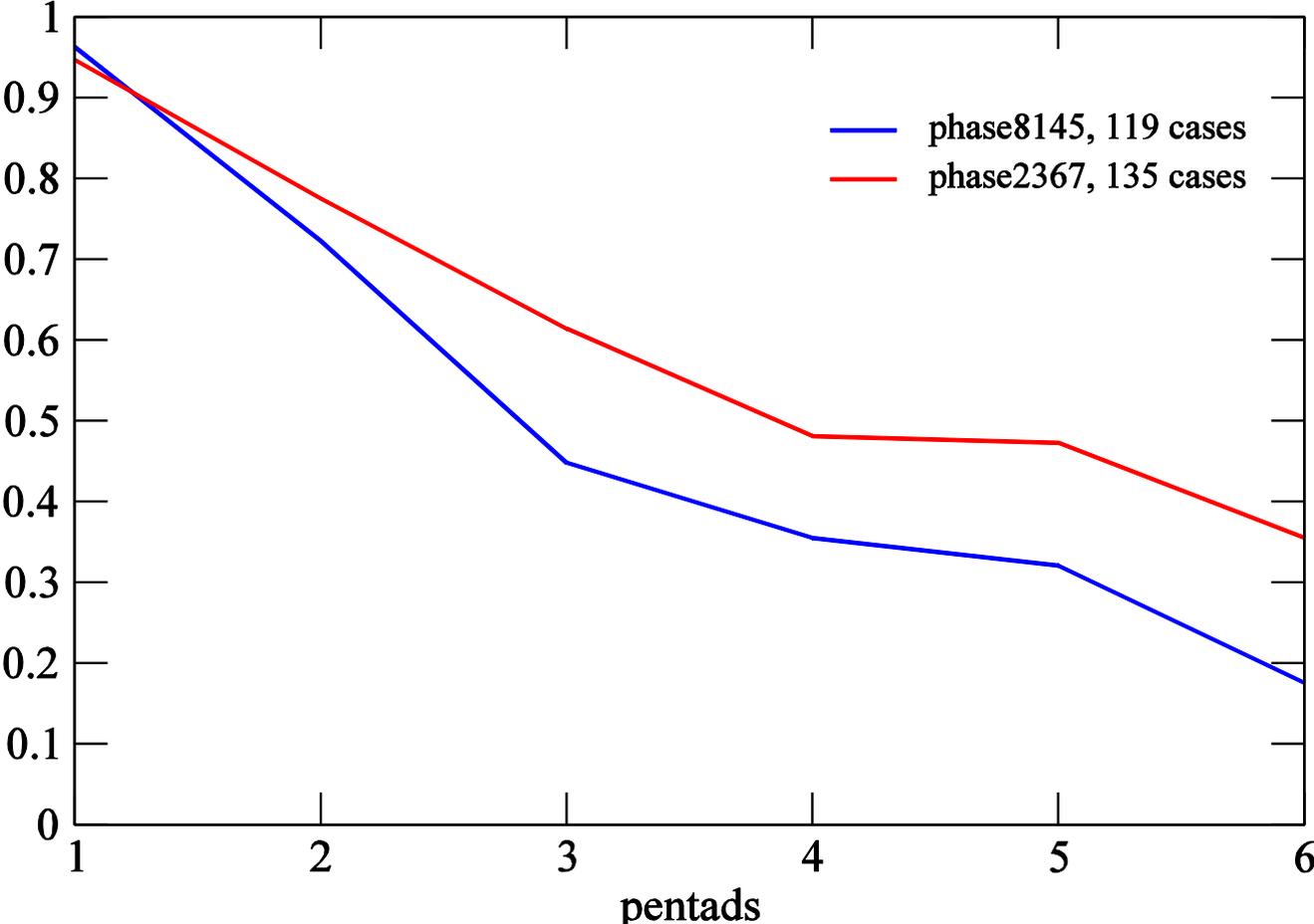


# **Impact of MJO-NAO interaction on subseasonal predictions**

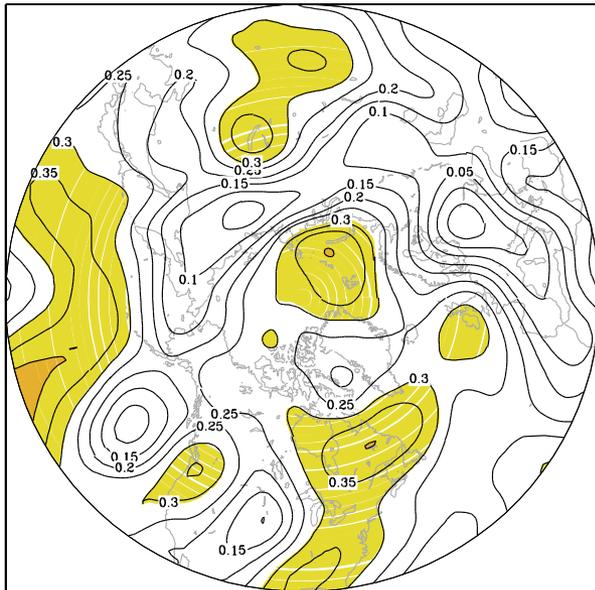
### NAO forecast skill



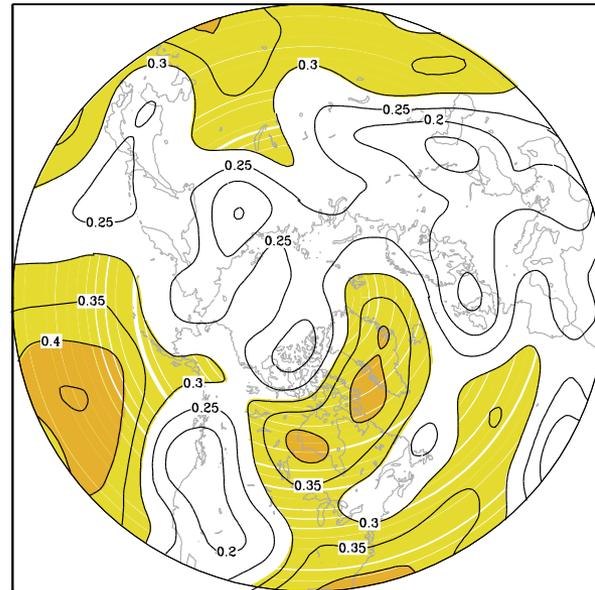
NAO forecast skill



a) weak MJO: Z500 skill

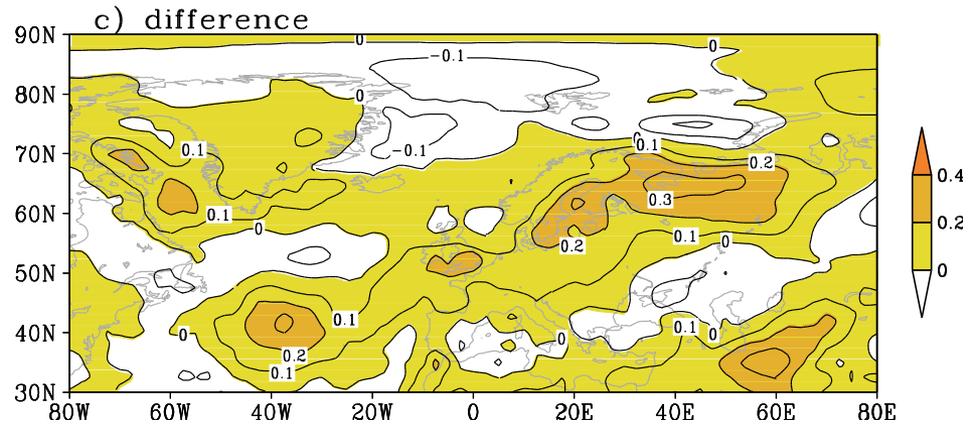
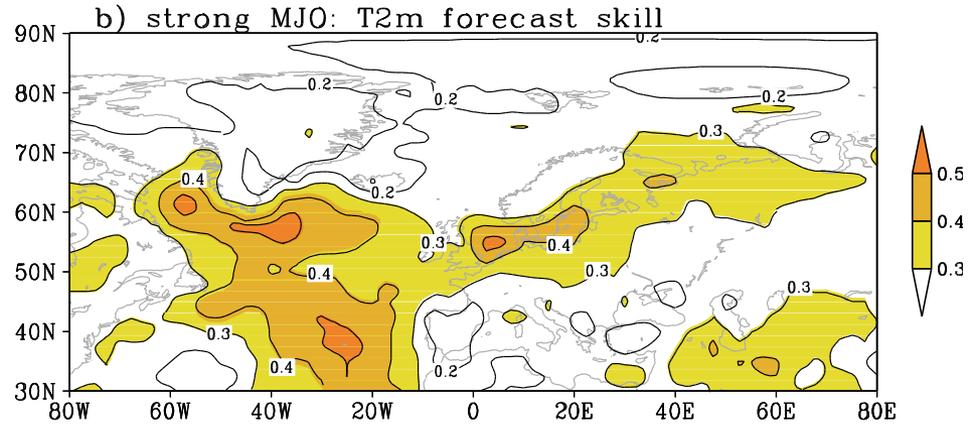
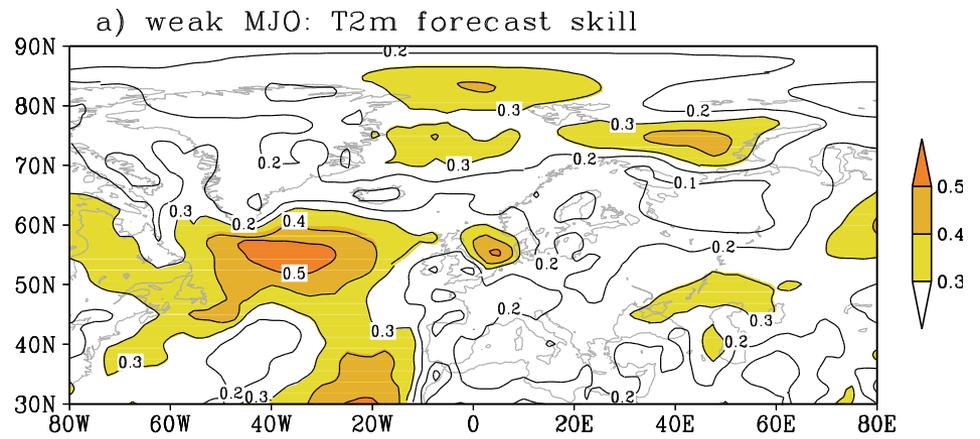


b) strong MJO: Z500 skill



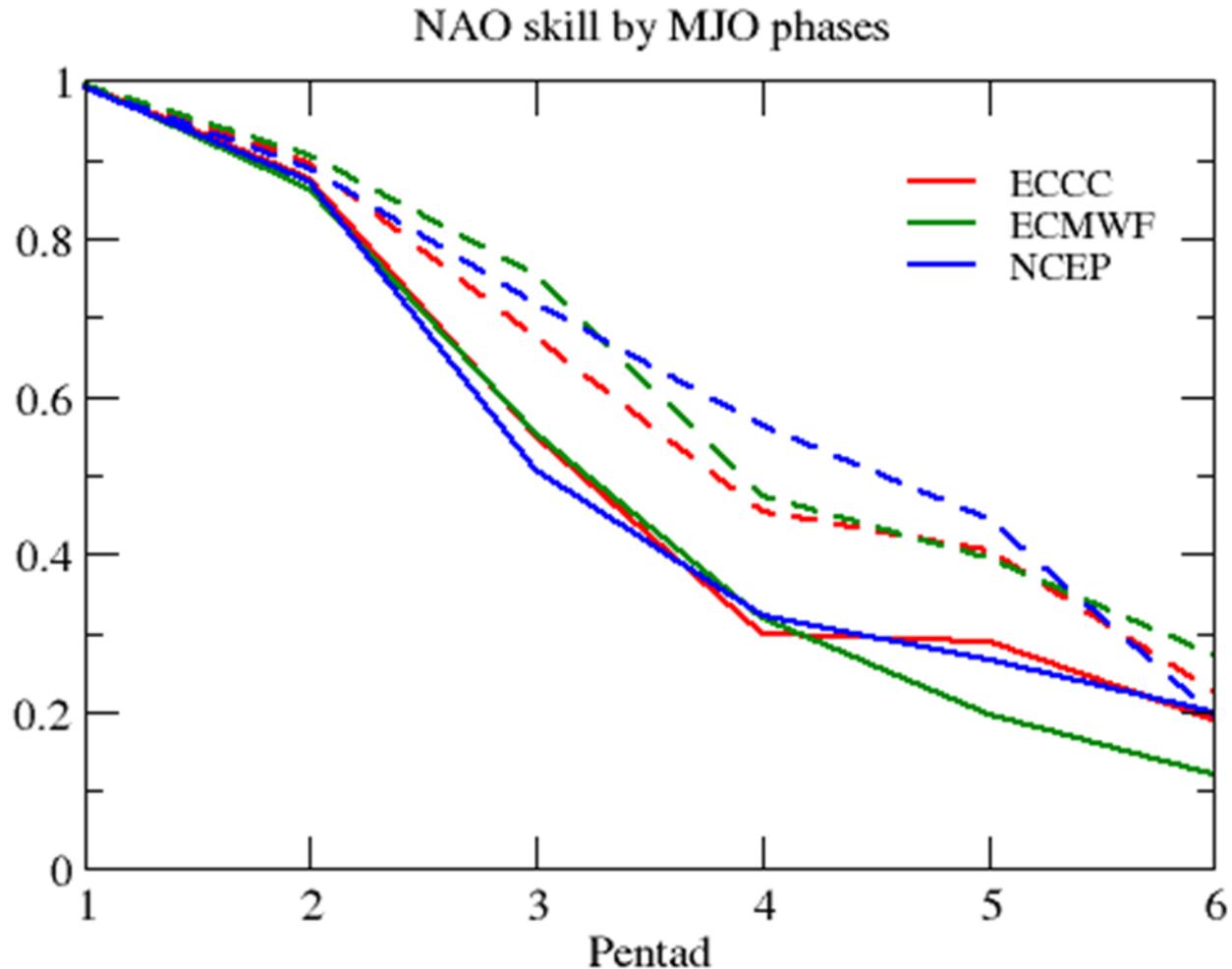
Correlation skill: averaged for pentads 3 and 4

Correlation skill: averaged for pentads 3 and 4





# S2S hindcast data

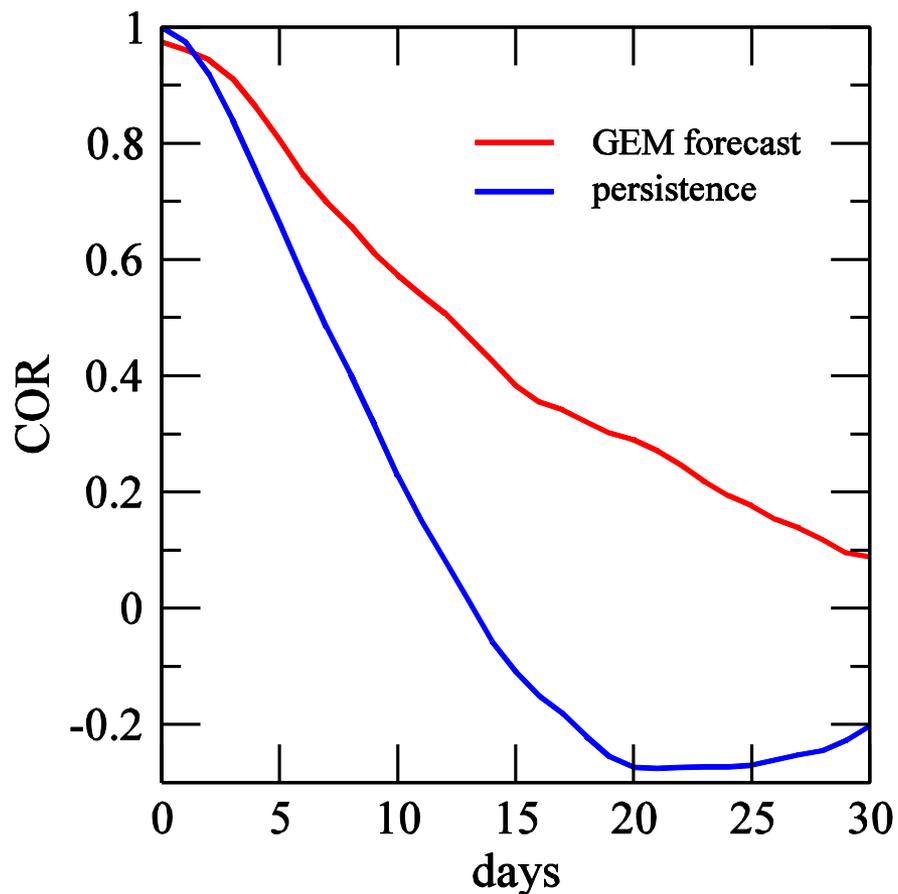


NAO forecast skill when the initial condition is in MJO phase 2367 (dashed) compared with MJO phases 1458 (solid).

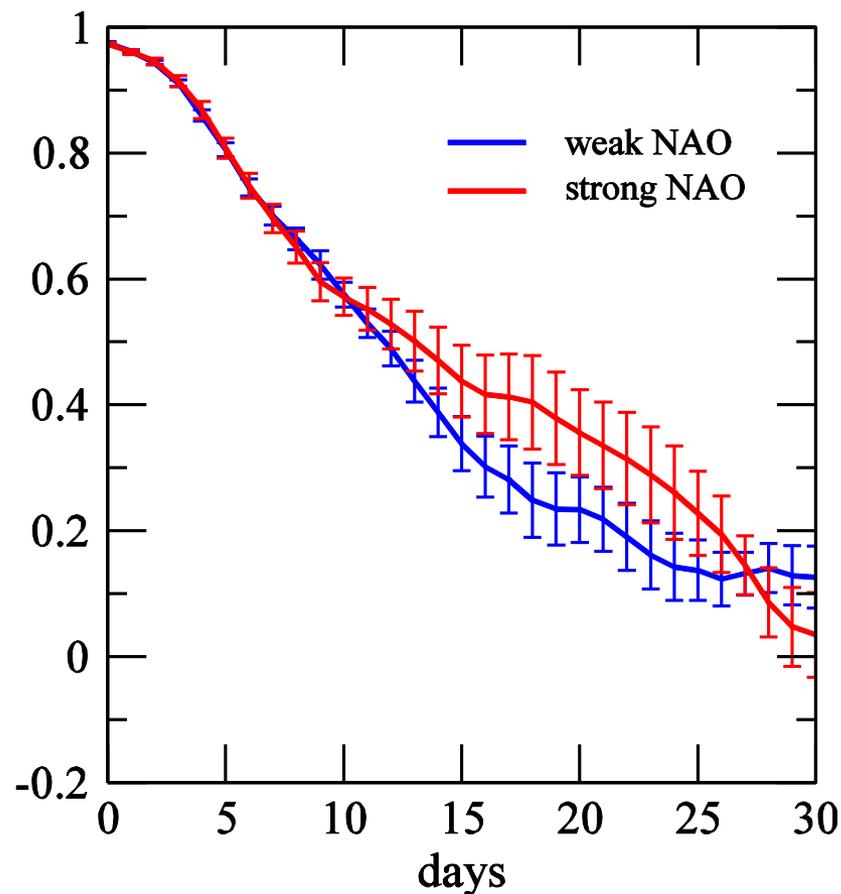
# **MJO forecast skill**

## **--- impact of the NAO**

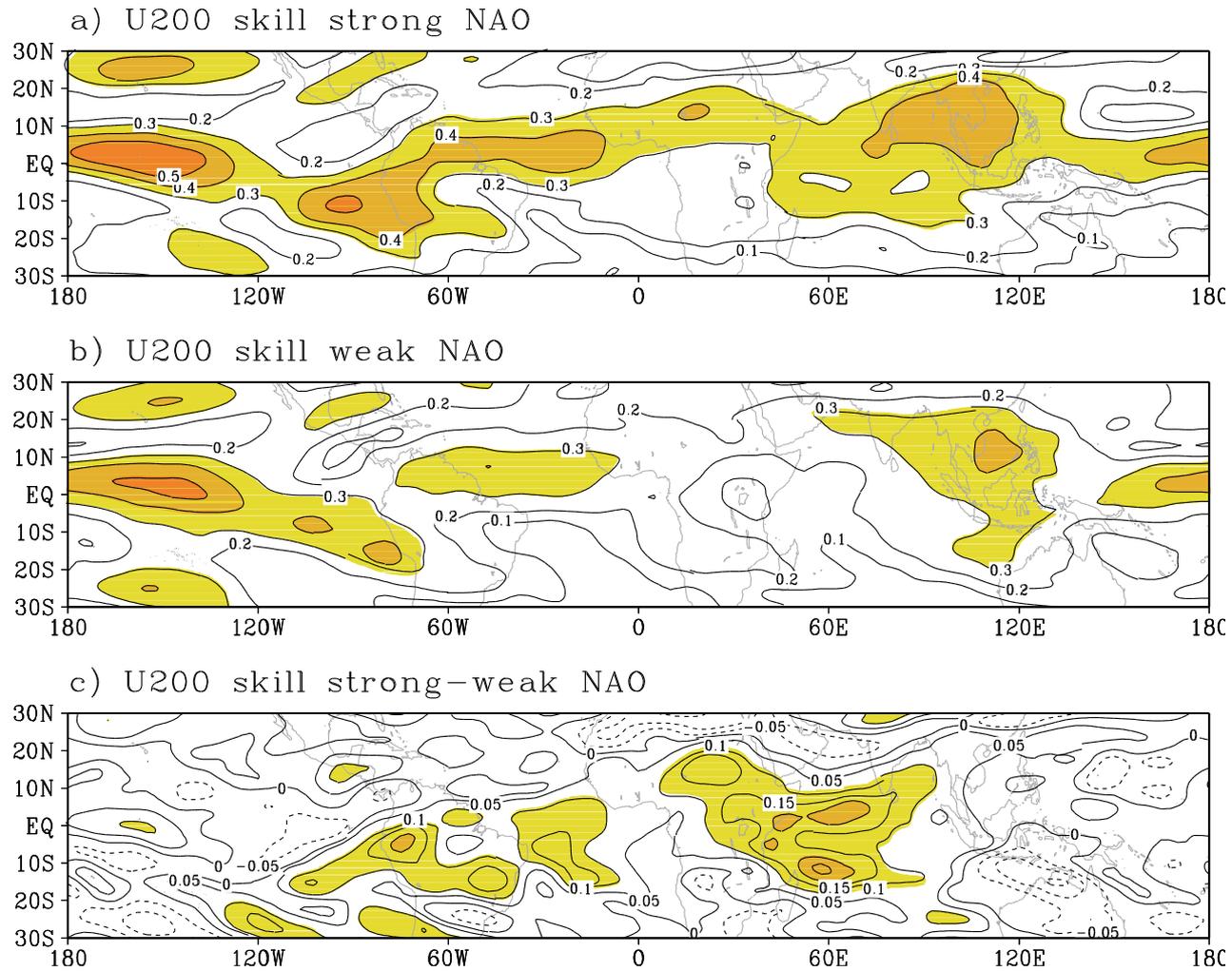
a) MJO forecast skill



b) MJO skill by initial NAO amplitude

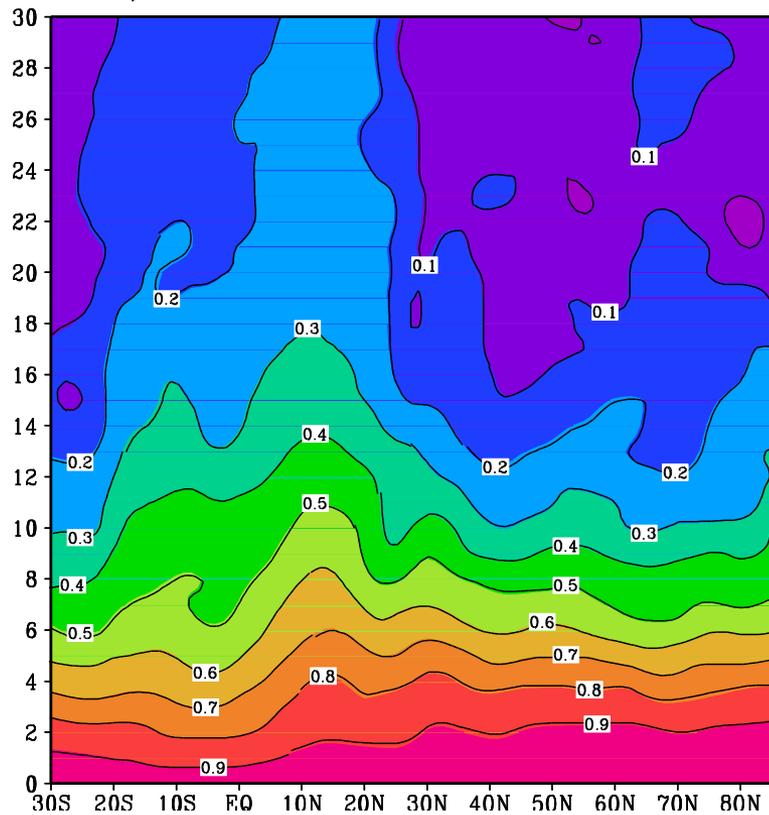


(Lin et al. GRL, 2010b)

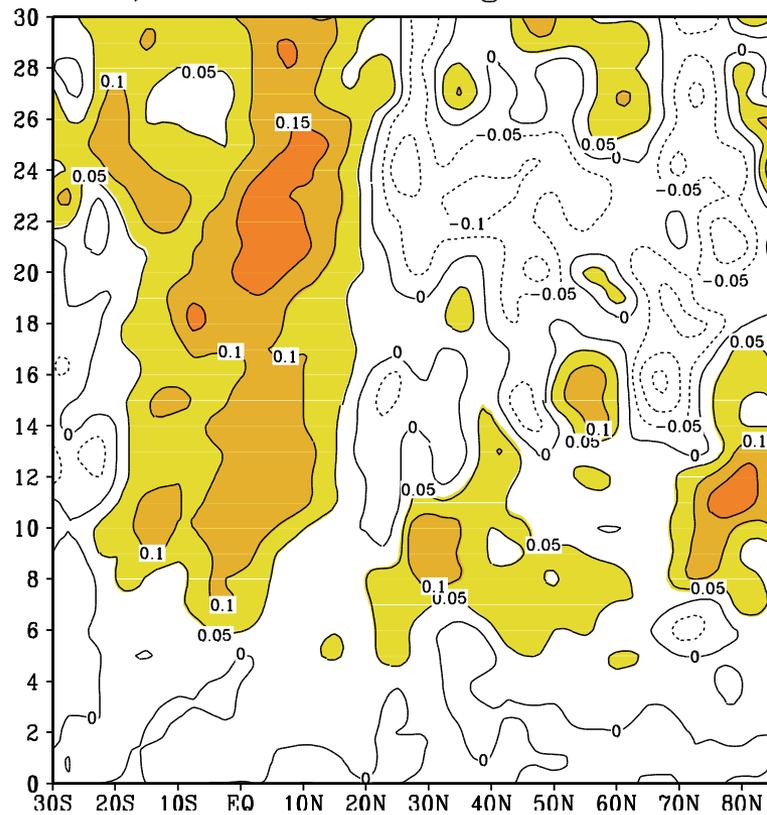


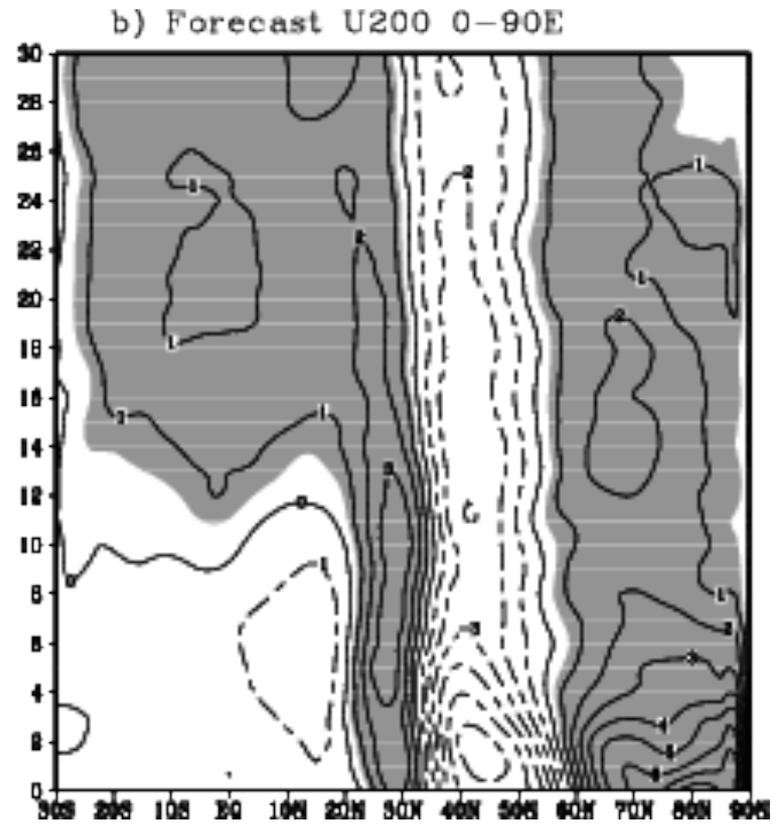
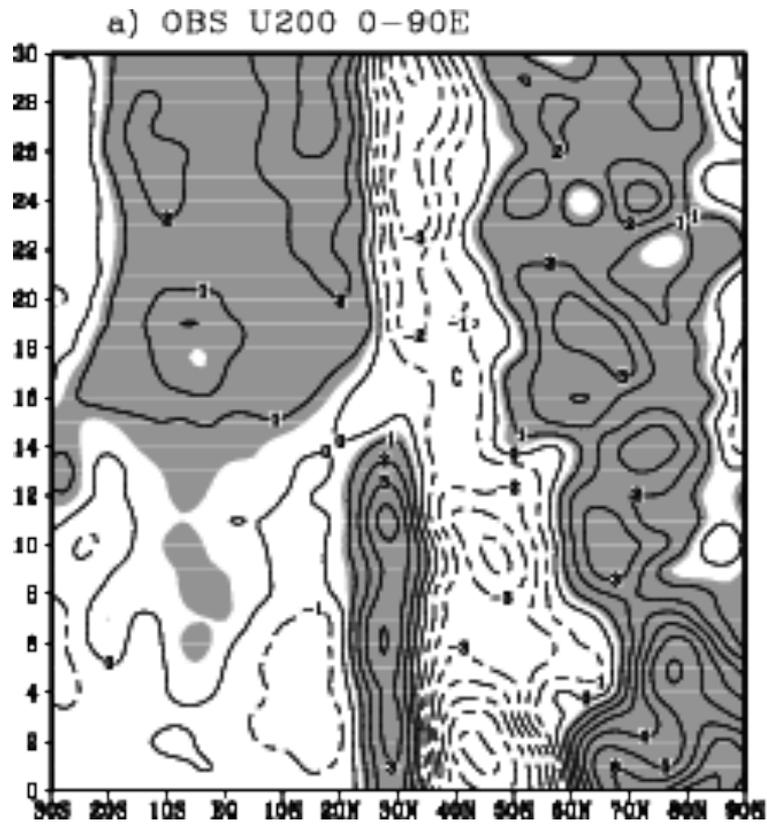
Skill averaged for days 15-25

a) U200 skill 0-90E



b) U200 skill strong-weak NAO





(Lin et al. GRL, 2010b)

# Summary

- Two-way interactions between the MJO and NAO
- Lagged association of North American SAT with MJO
- NAO intraseasonal forecast skill influenced by the MJO
- MJO forecast skill influenced by the NAO