Dr Andy Turner, NCAS-Climate & Department of Meteorology



SYSTEMATIC MODEL BIASES IN THE SOUTH ASIAN MONSOON



ICTP Targeted Training Activity: Monsoons in a changing climate 31 July-4 August 2017



OUTLINE

- CMIP5 performance for the Asian monsoon
- Arabian Sea as a source of bias
- Impact of coupling in CMIP5
- Parametrization uncertainty
- Role of resolution

Thanks to Fred Kucharski and J. Shukla for the invitation



ITCP TTA: Monsoons in a changing climate, August 2017 **MONSOON SIMULATION IN CMIP5**



MONSOON PRECIPITATION BIAS

- Large range of skill at simulating mean monsoon precipitation in CMIP5 & CMIP5 models
- Sperber *et al*. (2013)

Mean JJAS precipitation (left) and bias versus GPCP obs (right)









CMIP PERFORMANCE FOR ASIAN MONSOON



Mean JJAS precipitation (left) & bias versus GPCP obs (right) d) CMIP5 MMM i) CMIP5 MMM - GPCP 0.90 50N 301 30N 20N 10N Sperber *et al.* (2013, 13.5 Clim. Dyn.) CMIP3 MMM CMIP3 MMM - GPCP 0.86 301 301

- Large dry bias (India)/wet bias (WEIO)
- Incremental improvements only since CMIP3



MULTI-MODEL MEAN WINDS

• Weak Somali jet in CMIP5 and CMIP3



Mean JJAS 850hPa winds (left) and bias versus ERA-40 (right)



CMIP PERFORMANCE FOR ASIAN MONSOON

- Intimate link between monsoon circulation and precipitation biases
- Considerable effort still needed to improve coupled model performance
- Multi-model means outperform individuals

Sperber *et al.* (2013, *Clim. Dyn.*): Scatter diagrams of model pattern correlation skill for historical simulation of precip (y) and winds (x)



CMIP3 vs. CMIP5



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ARABIAN SEA AS SOURCE OF BIAS

SPRING TIME Reading ARABIAN SEA COLD BIASES

• CMIP5 models are too cold in winter and spring in northern Arabian Sea



Figure 1 | Arabian Sea SST bias. Northern Arabian Sea $(15^{\circ}N-25^{\circ}N; 60^{\circ}E-70^{\circ}E)$ spring season (March – May) SST bias (K) calculated for 44 CMIP5 models. The bias in historical all forcing simulations are computed with reference to fifty year (1951–2000) climatology of HadISST1.1. This figure is plotted using NCAR Command Language (NCL).

From Sandeep & Ajayamohan (2014) *Scientific Reports*

SST BIASES & LINK TO RAINFALL

 CMIP5 models with cold winter/spring Arabian Sea have a weakened seasonal cycle of rainfall, later onset



Levine, Turner, Marathayil & Martin (2013, *Clim. Dyn.*)

MODEL BIAS: ARABIAN SEA IN WINTER

- Links to Arabian Sea cold biases
- All DJF averages, i.e. winter monsoon





Marathayil/Turner/ Shaffrey (2013, *ERL*)

- Northern Arabian Sea SST cold biases linked to excess winter monsoon winds
- Excess LH flux lost to the atmosphere
- Yet air over north AS is too dry
- Advection of cold dry air across north coast of Arabian Sea

CONNECTION OF ARABIAN SEA COLD BIASES

- Composite of "coldest" minus "warmest" Arabian Sea models
- DJF average
- Linked series of coupled biases



☆ HiGEM × ukmo hadgem1 + ukmo hadcm3 ncar pcm1 × ncar⁻ccsm3 0 ☆ mpi echam5 ×miroc3 medres + miroc3 hires ★ ipsl cm4 ×inmcm3 0 + ingv echam4 ☆iap fqoals1 0 q giss model e r giss model e h 🖈 giss aom × afdl cm2 1 + afdl cm2 0 csiro mk3 5 csiro_mk3_0 cnrm_cm3 ★ cccma cgcm3 1 t63 × cccma cgcm3 1 t47 + bccr bcm2 0 observation



Marathayil/Turner/Shaffrey (2013, *ERL*)

IMPACT OF COLD BIASES ON MONSOON





60E

90E

120E

Comparison of monsoon rainfall in: > AGCM

- ≻ CGCM
- > AGCM forced with SSTs from the CGCM

From Levine & Turner (2012) *Climate Dynamics*

ISOLATE ROLE OF LOCAL SST





From Levine & Turner (2012) *Climate Dynamics*



ICTP TTA: Monsoons in a changing climate, August 2017 **IMPACT OF COUPLING IN CMIP5**

AMIP JJAS MEAN MONSOON



COURD COURD JJAS MEAN MONSOON University of Reading ICHEC



2 4 6 8 10 12 14 16 18 20 22

COUPLED-AGCM MEAN



CMIP5 ONSET PERFORMANCE

- Monsoon progression to the NW across much of the Asian monsoon domain
- "Wang & LinHo metric" for onset pentads
- Some models perform awfully
- CMIPx multi-model mean gets the right idea but systematically late



CGCM MINUS AMIP ONSET





























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CONVECTIVE PARAMETRIZATION





- Increasing the convective entrainment rate tends to improve ISV (e.g. Klingaman et al.; Hirons et al., 2012; Del Genio et al., 2012)
- Increasing convective entrainment globally decreases several biases while increasing others (Kim et al. 2011)
- Changing the entrainment rate has a dramatic impact on distribution of monsoon rainfall, but doesn't solve the fundamental bias

Bush *et al.* (2014,



ICTP TTA: Monsoons in a changing climate, August 2017 **ROLE OF RESOLUTION**

ROLE OF RESOLUTION



- Potential for better resolved processes
- Less to be achieved by sub-gridscale parametrizations



Johnson *et al*. (2016, *Clim. Dyn*.)

IMPACT OF RESOLUTION ON RAINFALL

 For the Asian monsoon region as a whole, increasing resolution at these scales doesn't appear to improve simulation





P (mm day")

Johnson *et al*. (2016, *Clim. Dyn*.)





10S 20S University of **Reading**





RESOLUTION & DEPRESSIONS

- In these simulations, resolution does not seem to change number of LPS (not shown)
- Rainfall associated with LPS does increase

Johnson *et al*. (2016, *Clim. Dyn*.)



Fig. 7 Climatological, ensemble averaged JJAS precipitation attributed to monsoon LPS in the a N96 configuration of the MetUM on its native grid and b N512 configuration of the MetUM interpolated to the N96 grid. c Difference between b and a. Only points significant by a Mann–Whitney rank sum test are shown. The *colour scale* ranges to ± 1.5 mm day⁻¹, while the scale on Fig. 3, which shows the total JJAS precipitation change with resolution, ranges to ± 5 mm day⁻¹

RESOLUTION & OROGRAPHY



- 11-13°N mean
- Clear increase in rainfall peaks associated with steep orography, but still failing to match
 - match observations

Johnson *et al*. (2016, *Clim. Dyn*.)





RESOLUTION & MONSOON FLOW



- Clear and systematic impact of resolution in increasing speed of Somali jet
- But N512 resolution still slower than reanalysis



RESOLUTION & MONSOON FLOW

- 10-15°N east/ west crosssection
- Meridional flow shown shaded
- East African Highlands better resolved by N21(resolution
- High resolution improves both zonal flow and meridional crossequatorial flow Johnson et al. (2016, Clim. Dyn.)





SUMMARY & OUTLOOK

- CMIP models capture the basic features of monsoon, including NW onset propagation
- Large dry biases coupled to weak monsoon winds
- Coupled models seem to perform worse than AMIP models due to cold SST biases developing in the Arabian Sea
- Distribution of tropical (& Asian monsoon) rainfall is sensitive to the *detail* of convective parametrization
- Resolution improves some features of the monsoon, but at typical GCM scales it does not solve the bias problem



THANK YOU!

See:

- Sperber et al. (2013) Climate Dynamics
- Marathayil et al. (2013) Environ. Res. Letts.
- Levine and Turner (2012) Climate Dynamics
- Levine et al. (2013) Climate Dynamics
- Johnson et al. (2016) Climate Dynamics
- Bush et al. (2015) QJRMS

Spatial variations in the monsoon



Overall INCOMPASS flight strategy:

- To sample spatial contrasts across northern India in the premonsoon and as the onset progresses
- 2 To sample contrasts across southern India in the mature monsoon



INCOMPASS project part of the NERC/MOES Monsoons Programme 2015-2018





Surface flux observations



INCOMPASS project part of the NERC/MOES Monsoons Programme 2015-2018



(Koster et al., Science, 2004)



- Huge area equipped for irrigation in northern India
- Evidence in models for strong coupling between land and atmosphere in this region
- Contrasts between wet and dry soils

Despite all these factors, measurements of the land and its interaction with the atmosphere are sorely lacking



NDIAN INSTITUTE OF SCIENCE angalore, India ব্যীয থিয়ান অধ্যযান







IIT-Kanpur supersite (~85km from LKO)



Flux tower: permanent installation; surface flux data sent via mobile network to UK Lidar ceilometer: permanent installation; test data have successfully tracked height of cloud base Microwave radiometer: permanent, has begun testing

Radiosonde receiving station: intensive observations during July 2016

INCOMPASS project

NERC/MoES Monsoons Programme 2015–2018