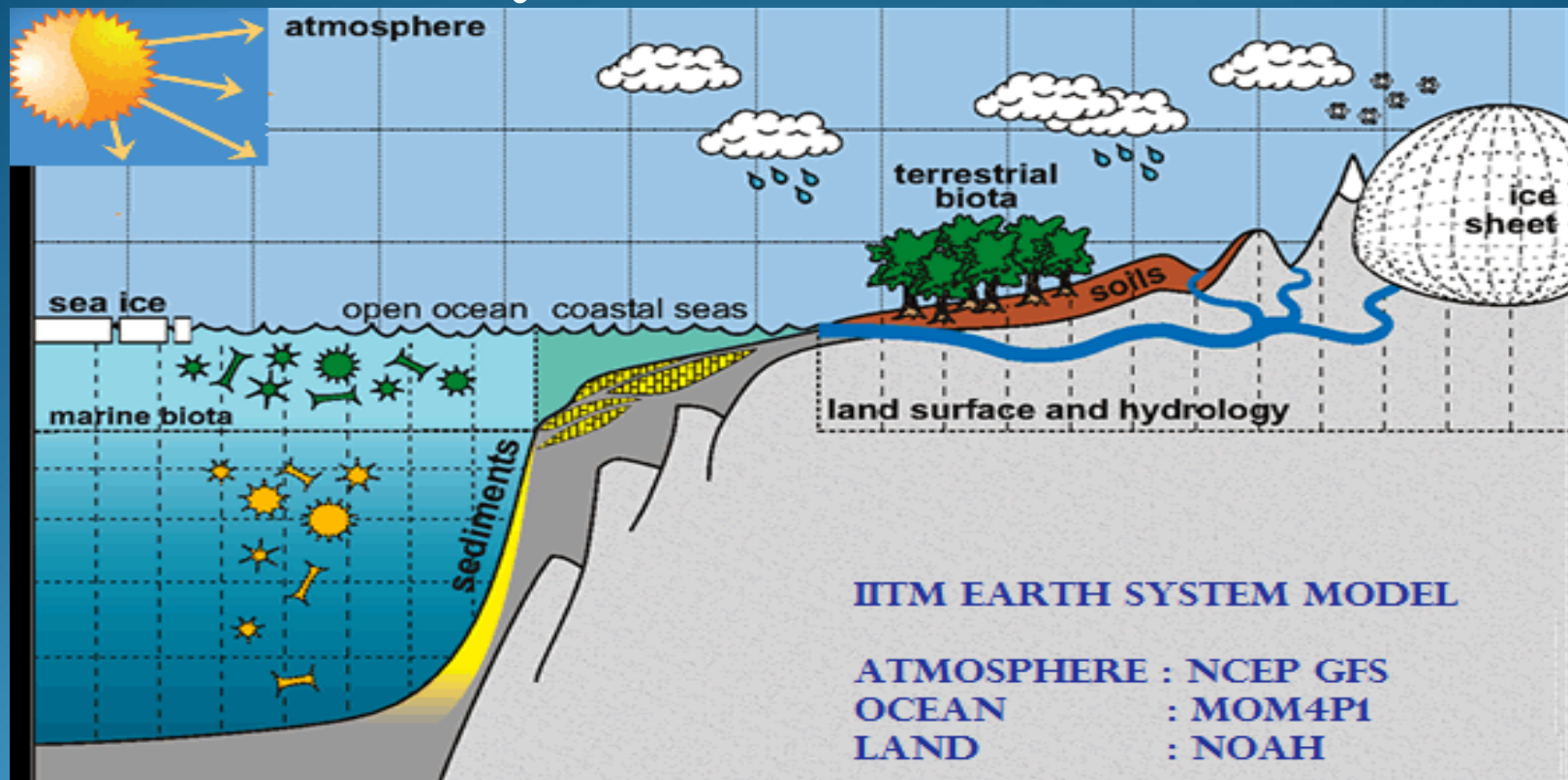
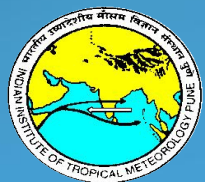


Monsoon Simulation in IITM ESM : Present Day and Future Scenario



Swapna Panickal

Centre for Climate Change Research
Indian Institute of Tropical Meteorology (IITM)



Background

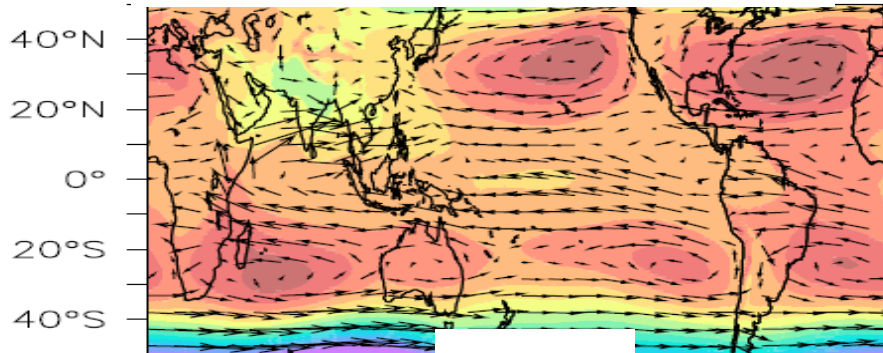
- **Reliable predictions of Asian monsoon rainfall to climate change are extremely important; lives of nearly half of the world's population depend on monsoon for food and energy security**
- **Understanding how the monsoon will change in the future is a fundamental challenge for climate science**
- **The Coupled Modeling Inter-comparison Project coordinated by the World Climate Research Programme (WCRP) form the basis of the climate projections in the Intergovernmental Panel on Climate Change (IPCC) Assessment Reports**
- **To address the future monsoon projections to climate change using the CMIP6 simulations of IITM-ESM, first from India.**

Outline

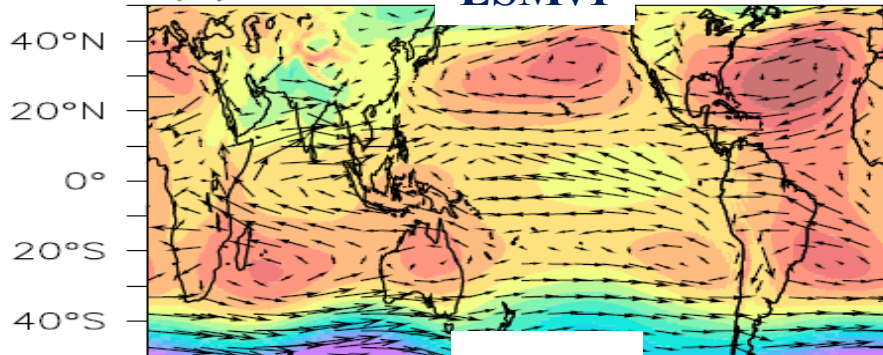
- **Present day Mean Climate from IITM-ESM Simulation**
- **Mean Features during Boreal Summer Monsoon Season**
- **Land-Atmosphere-Ocean Feedbacks and South Asian Monsoon**
- **ENSO and IOD Teleconnections on South Asian Monsoon Variability**
- **Impact of doubling CO₂ on Indian Summer Monsoon**

Large-scale Atmospheric Circulation

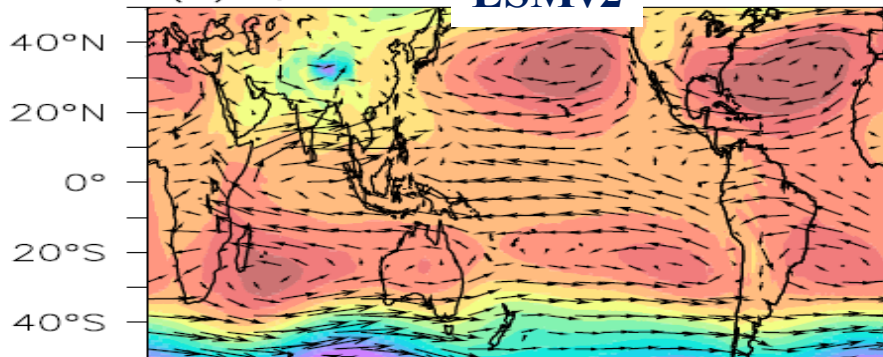
GP & Wind, Obs, JJAS, 850 hPa



(b) ESMv1 ESMv1

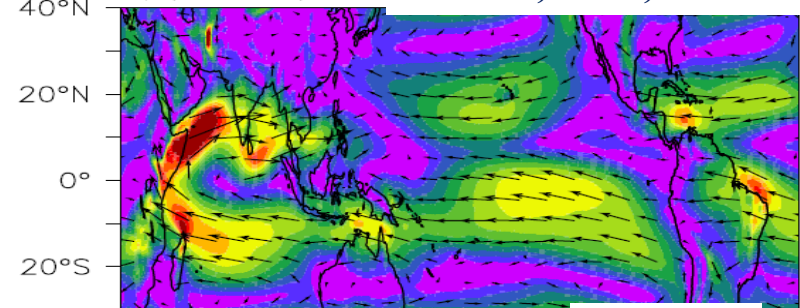


(c) ESMv2 ESMv2

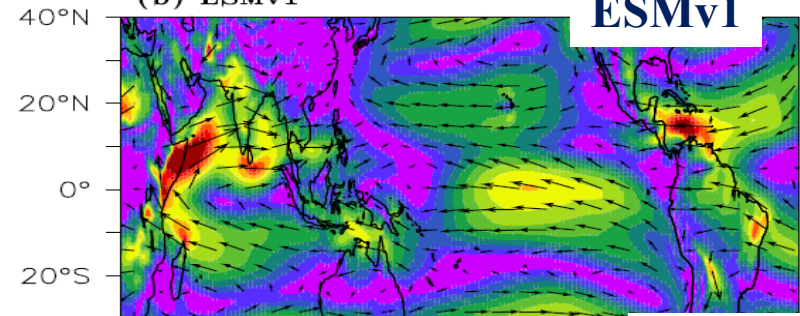


0° 100°E 160°W 60°W C

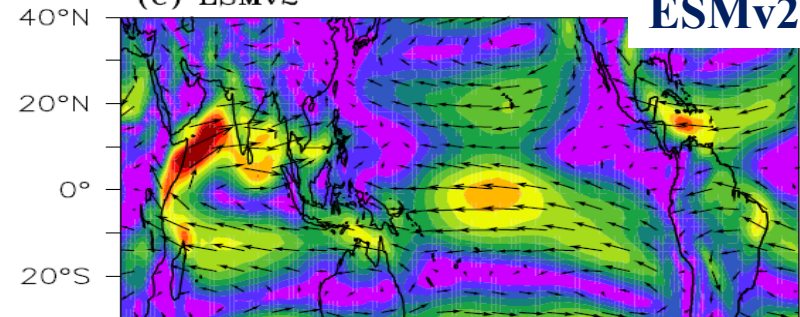
(a) Wind (Obs, Wind, Obs, JJAS



(b) ESMv1 ESMv1



(c) ESMv2 ESMv2

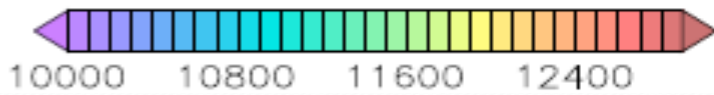
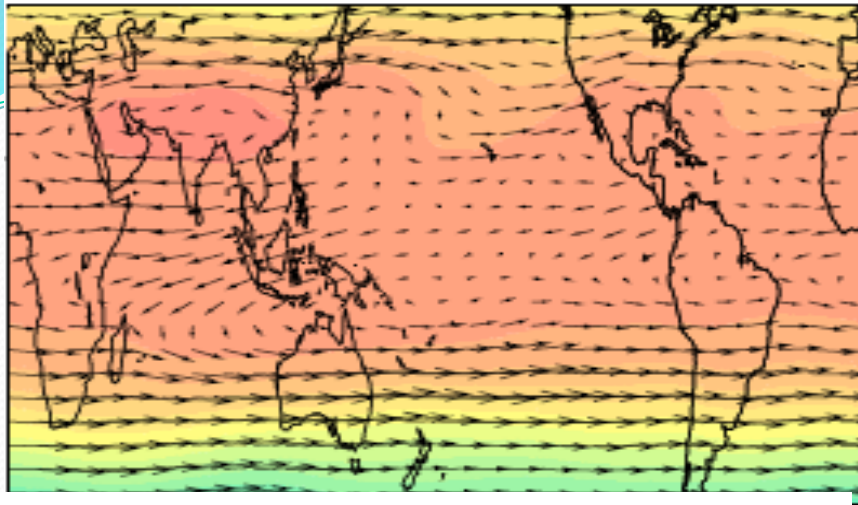


30°E 90°E 150°E 150°W 90°W 30°W

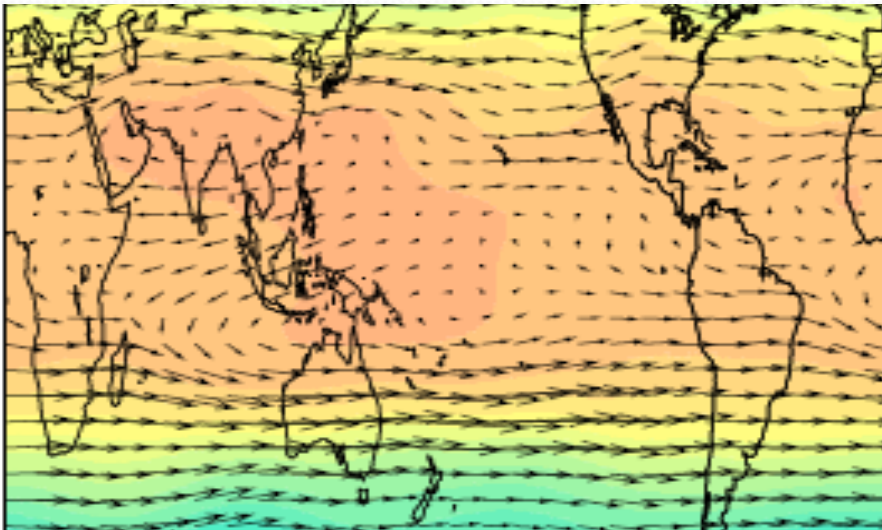


Winds & Geopotential Height: 200 hPa, JJAS

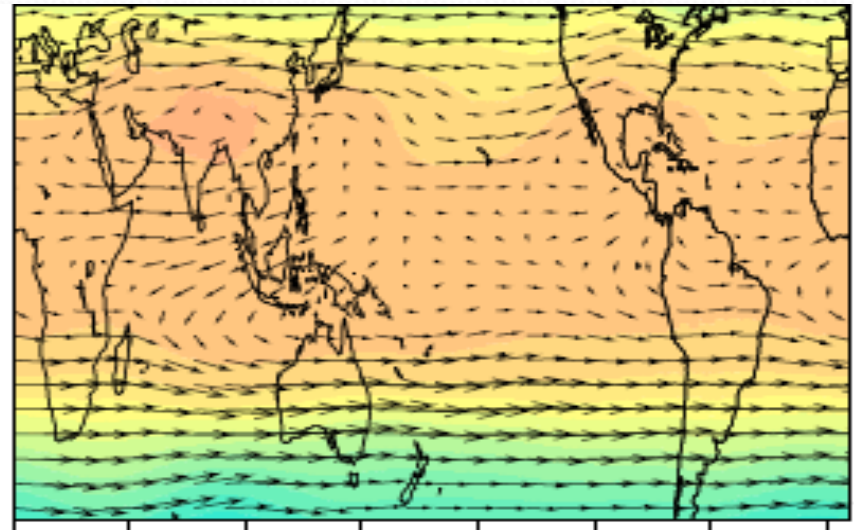
ERA-Interim



ESM-v1

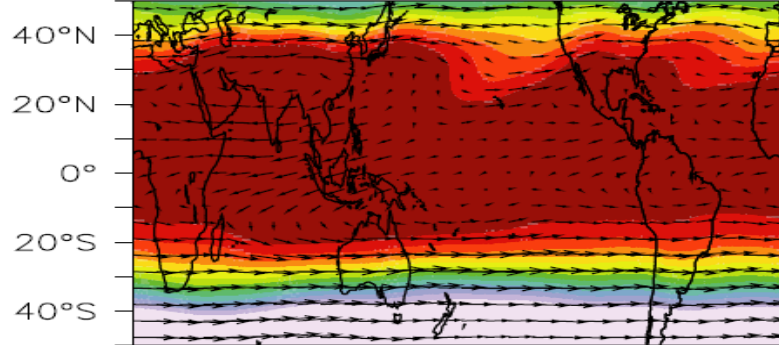


ESM-v2

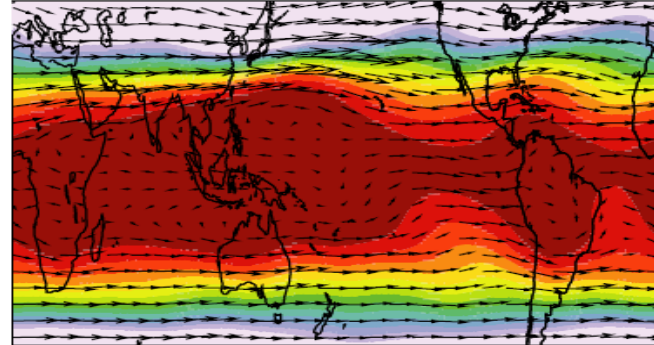


Winds & Geopotential Height: 200 hPa

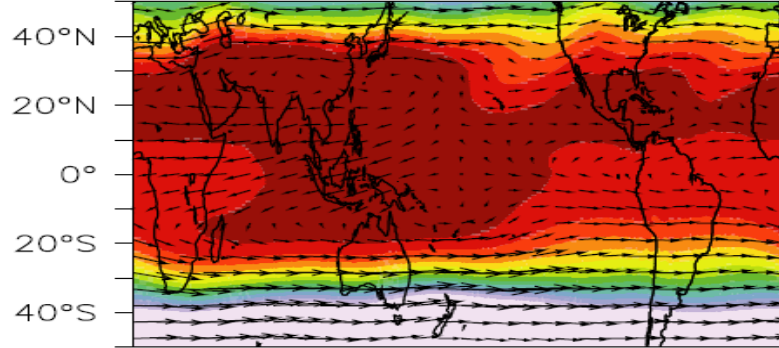
(a) GP & Winds (Obs, JJAS, 200 hPa)



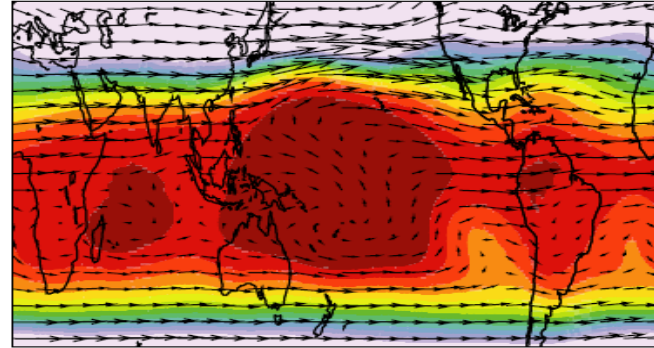
(d) Obs (DJF)



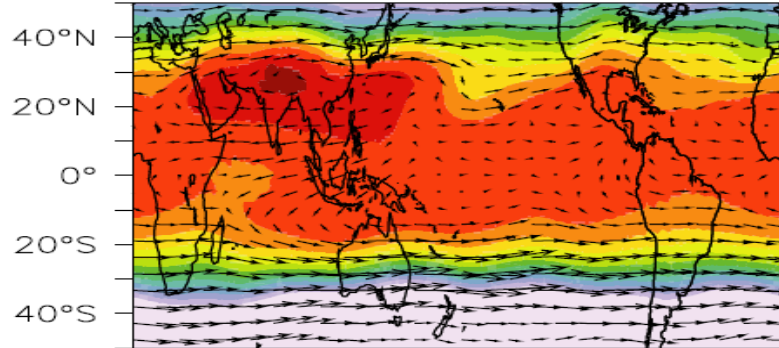
(b) AGCM



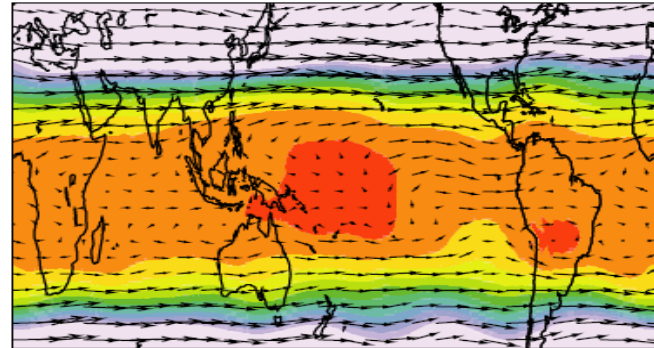
(e) AGCM



(c) ESMv2



(f) ESMv2

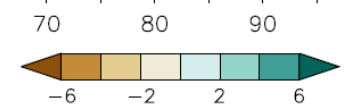
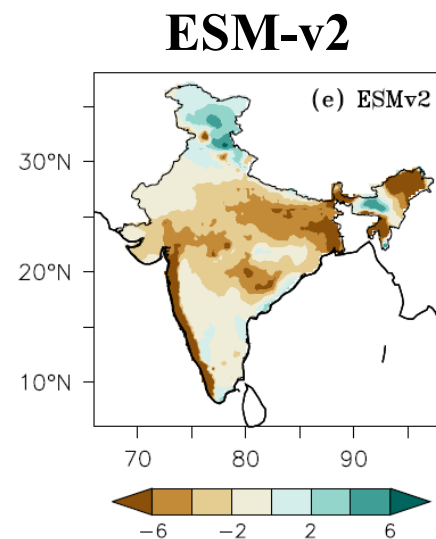
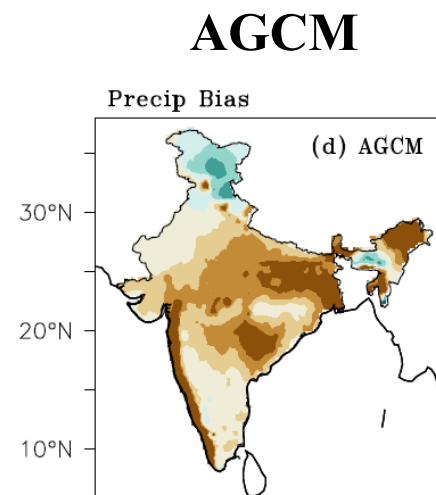
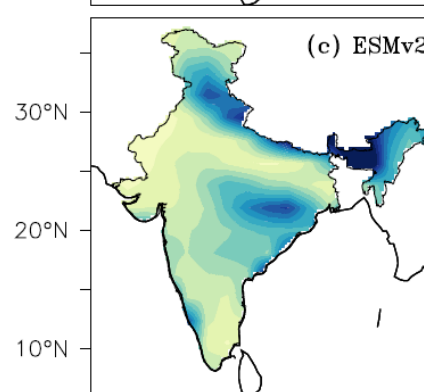
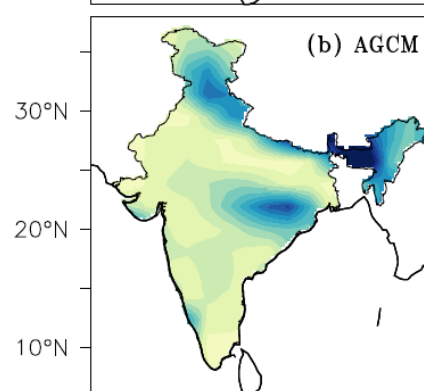
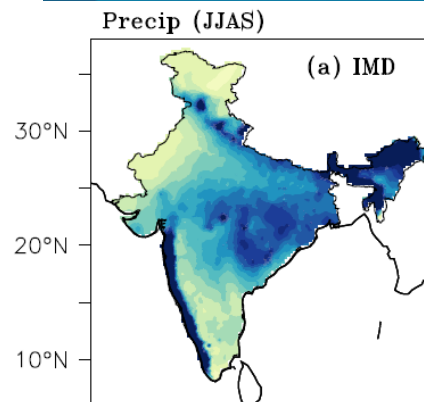
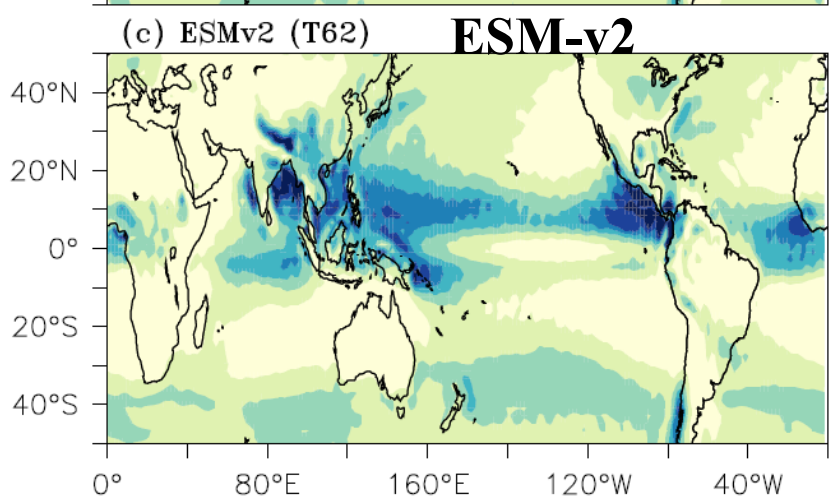
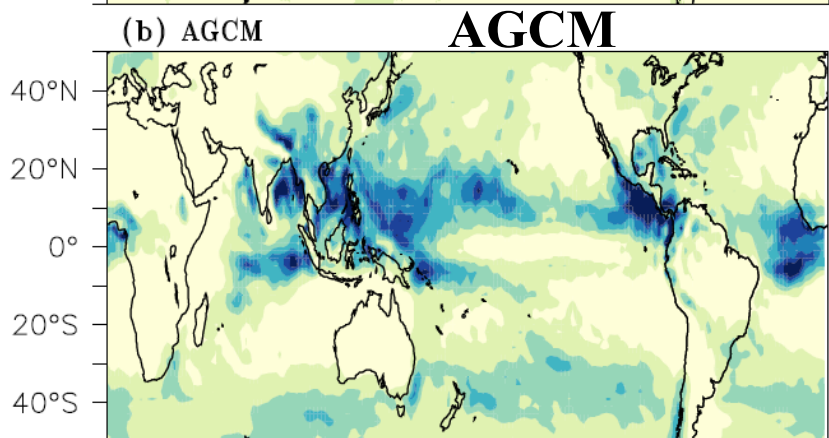
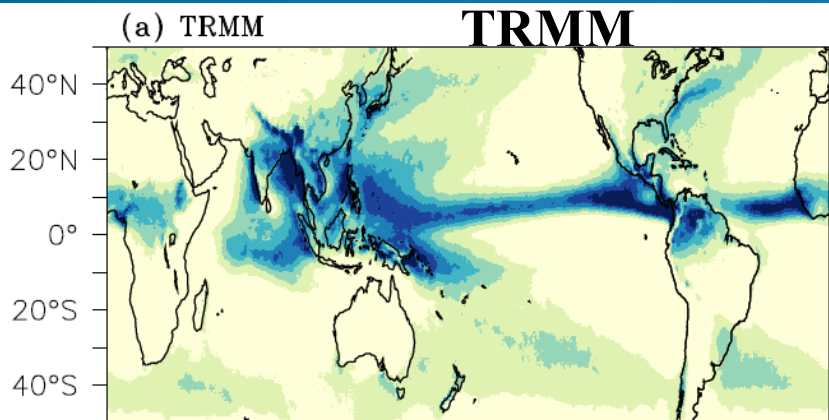


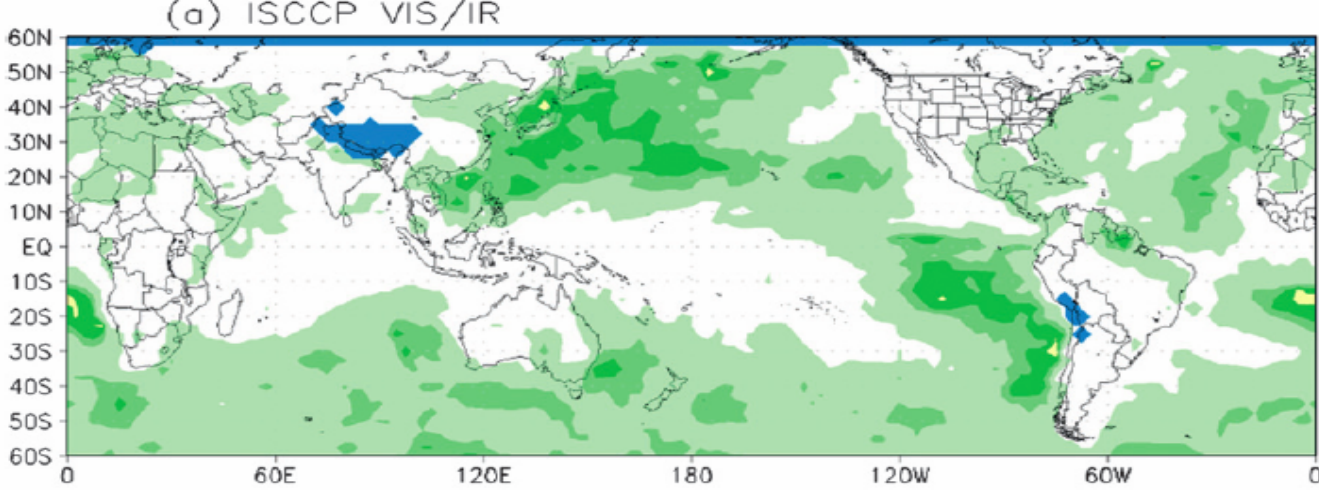
ERA-Interim

AGCM

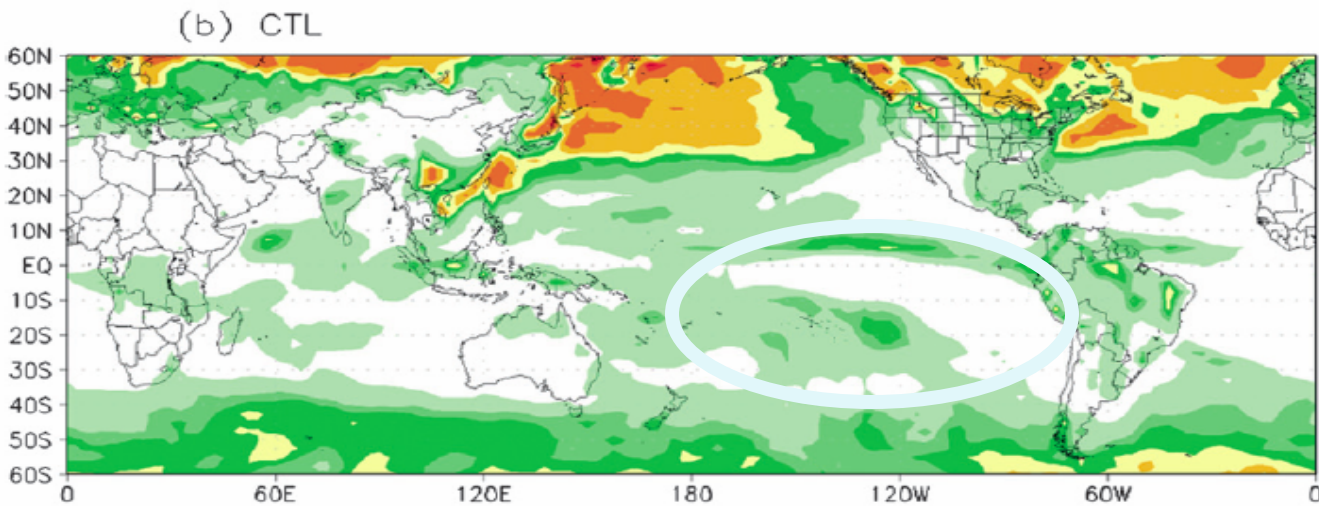
ESM-v2







Monthly mean low cloud cover (%) for January 2003 from ISCCP (Rossow and Schiffer, 1991) VIS/IR satellite observations (blue color indicates 'no data' available).



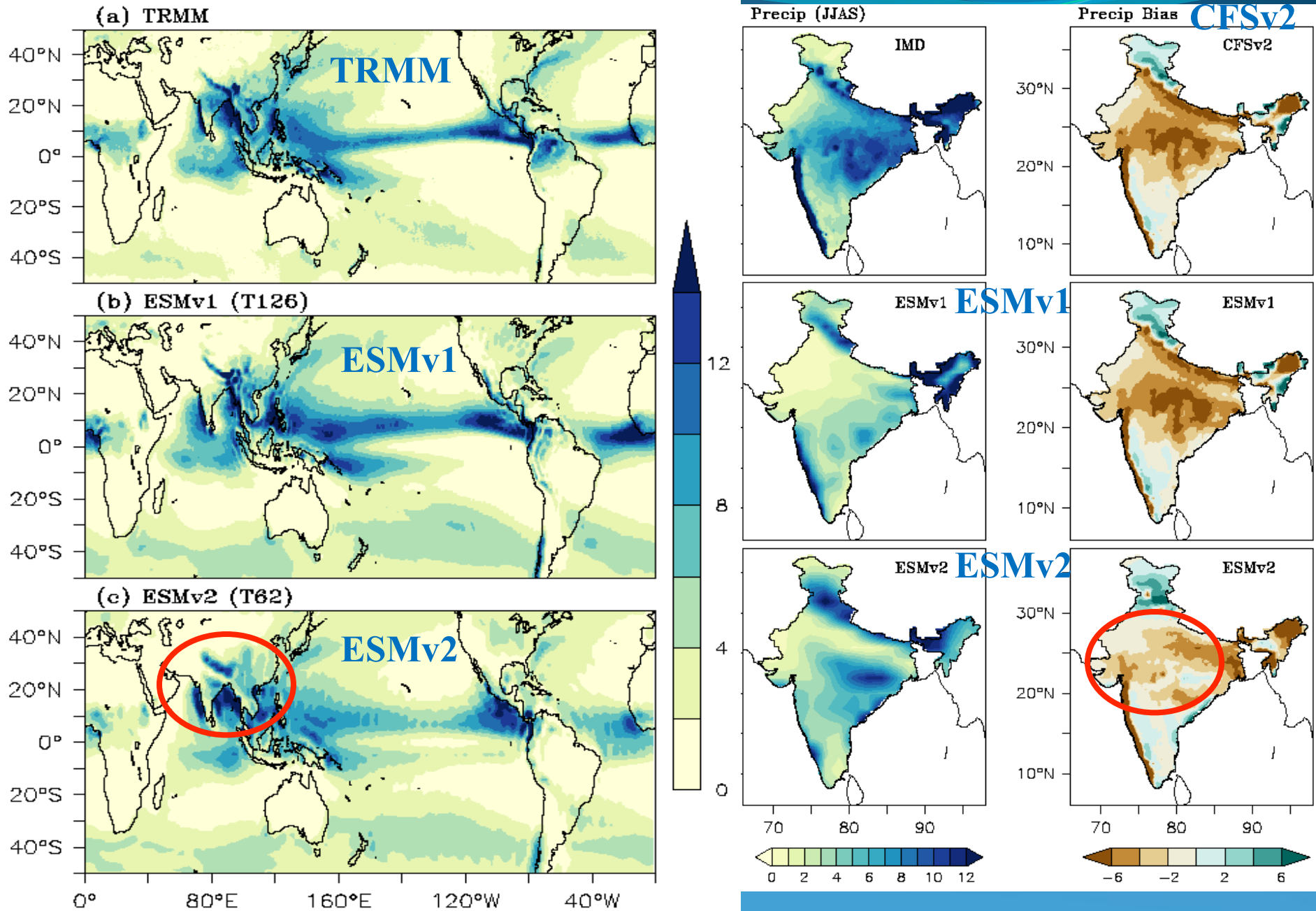
Control simulation using the old shallow convection Scheme of NCEP GFS

Han and Pan, 2011

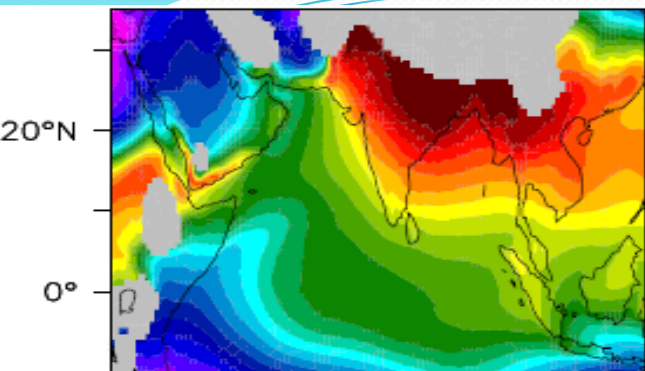


Long-standing problems in NCEP GFS: Systematic underestimation of stratocumulus clouds in the eastern Pacific and Atlantic Oceans; and the frequent occurrence of unrealistic excessive heavy precipitation, the so-called grid-point storms

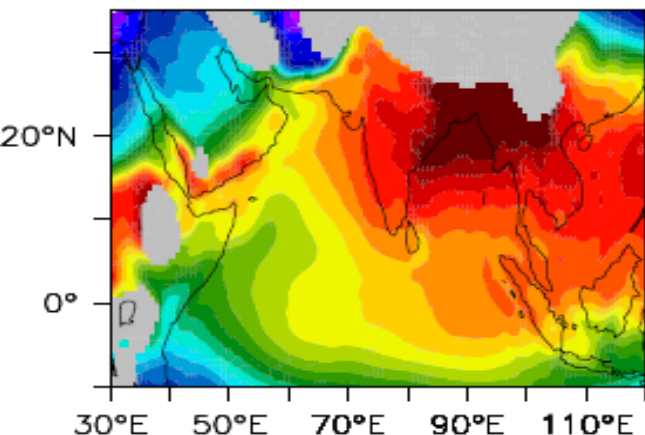
Boreal summer monsoon (JJAS) precipitation (mm day⁻¹)



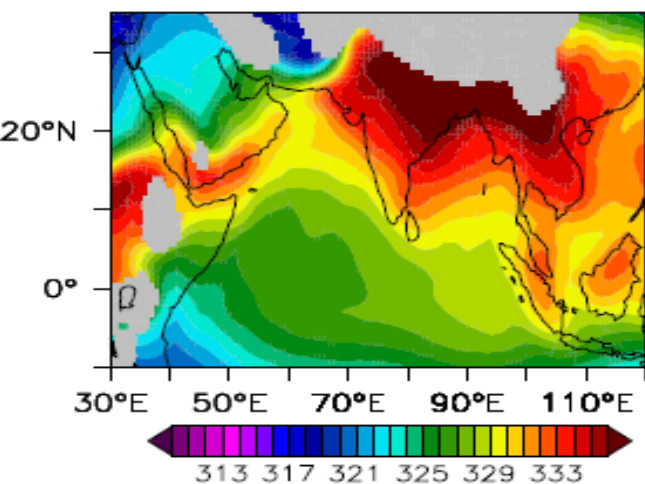
Moist Static Energy



ERA
Interim

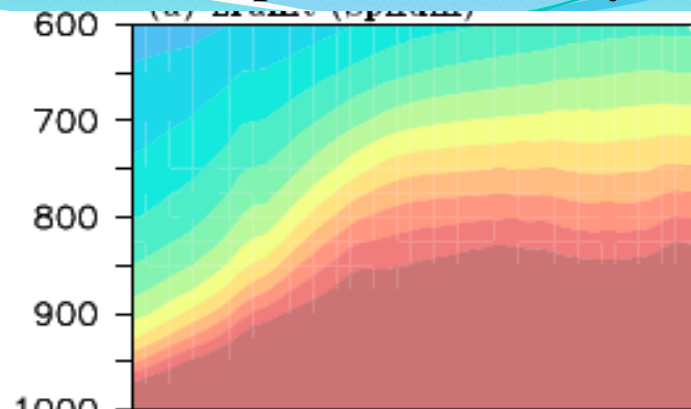


ESMv1

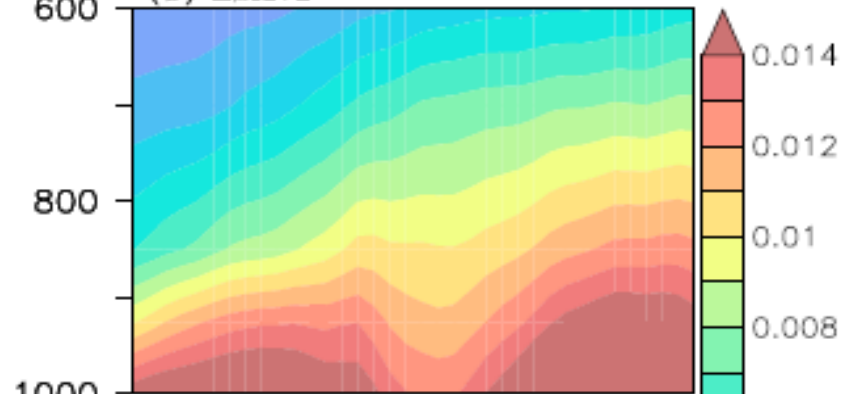


ESMv2

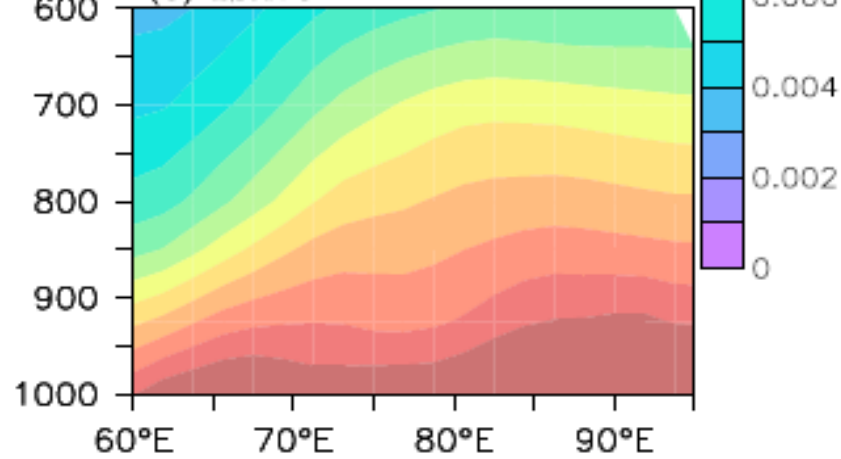
Specific Humidity



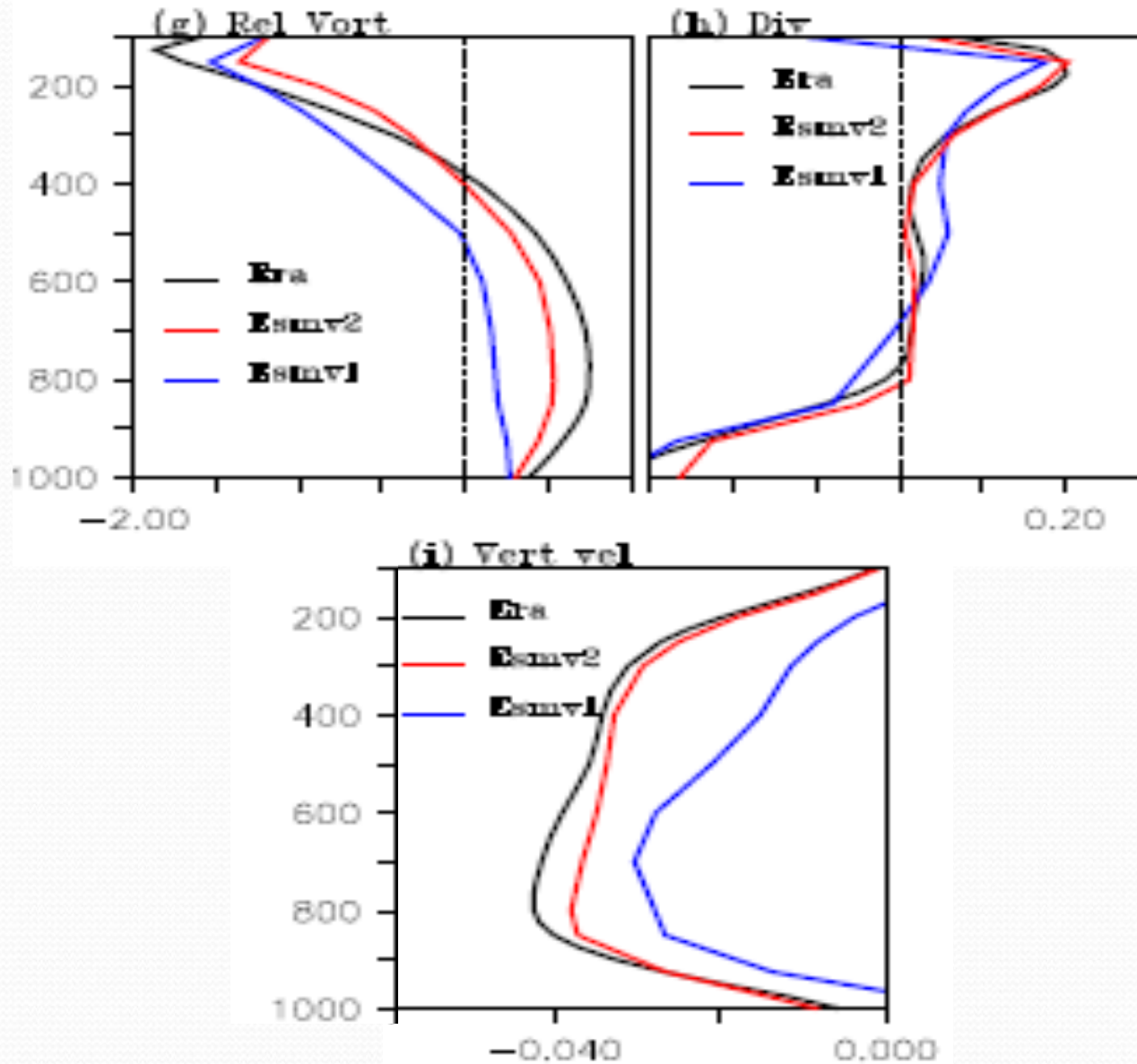
(b) ESMv1



(c) ESMv2

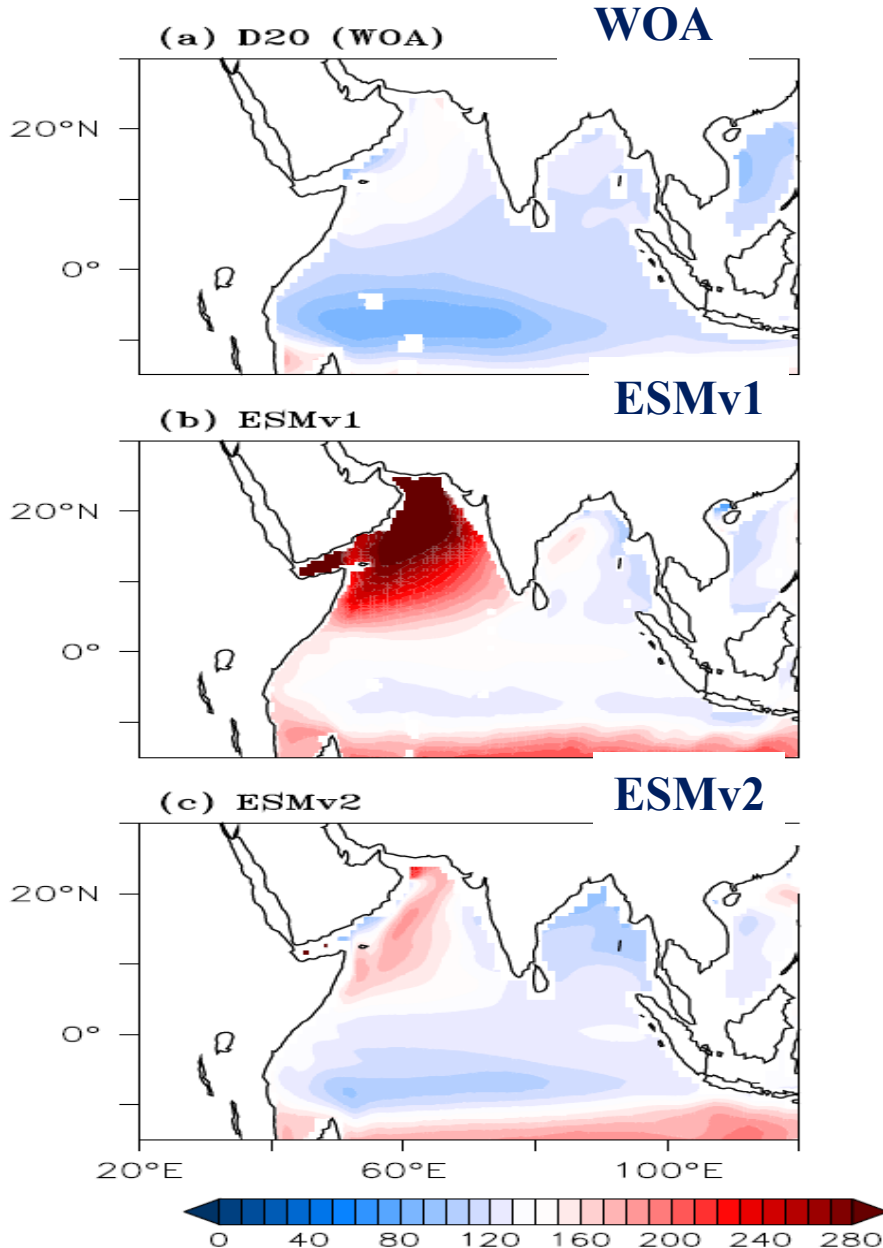


Moist processes over South Asian monsoon region



Indian Ocean-Monsoon Coupled Feedbacks

Indian Ocean-Monsoon Coupled Interaction



Improved cross-equatorial flow

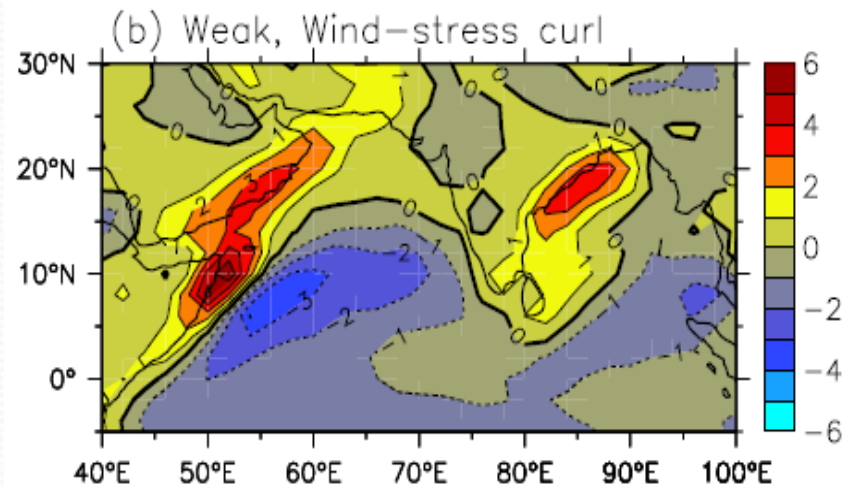
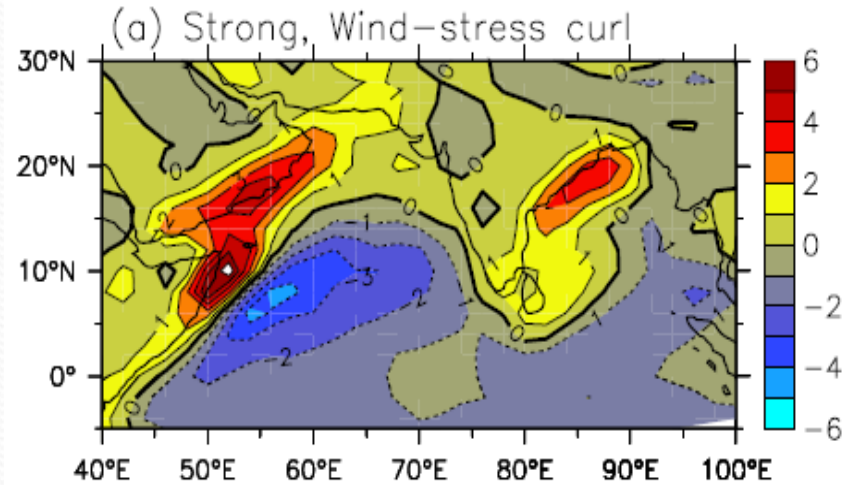
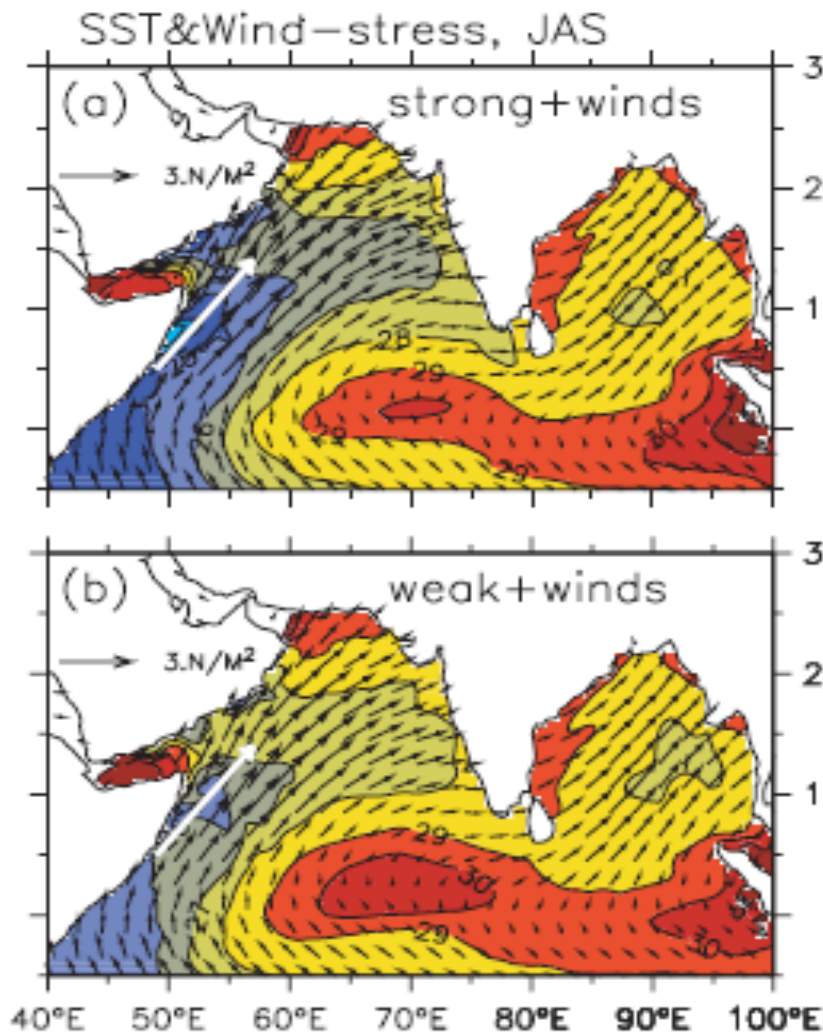


**Improved Ekman Upwelling,
shoaling of the SWIO thermocline
and thermocline dome**



**The shallow SW IO thermocline
dome warm the SST through
westward propagating Rossby waves,
generate zonal SST Gradient and
influence the IOD and IOBM**

Indian Ocean-Monsoon Coupled Interaction

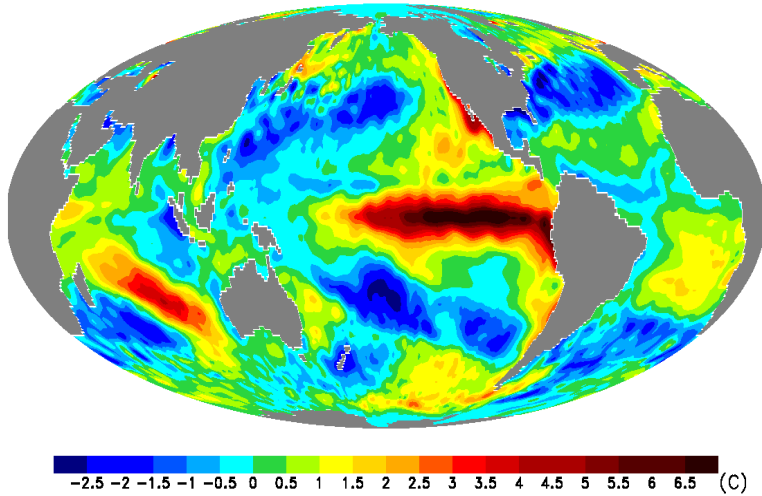


Arabian Sea response to monsoon variability (Murtugudde et al., 2007)

1997-98: Strongest El Niño ever recorded!

Sea Surface Temperature

Dec 1997 minus Dec 1998



In January 1998 (top right) the 1997-1998 El Niño event was at its height. Because of the weakness of the trade winds at this time, the upwelling of nutrient-rich water was suppressed in the equatorial Pacific. The absence of a green band along the equator in this image is indicative of relatively low chlorophyll concentrations there.

By July 1998 (bottom right) the trade winds had strengthened and equatorial upwelling had resumed giving rise to widespread phytoplankton blooms in the equatorial belt

(Ref: Wallace and Hobbs, 2006)

SeaWiFS Captures El Niño - La Niña Transitions in the Equatorial Pacific

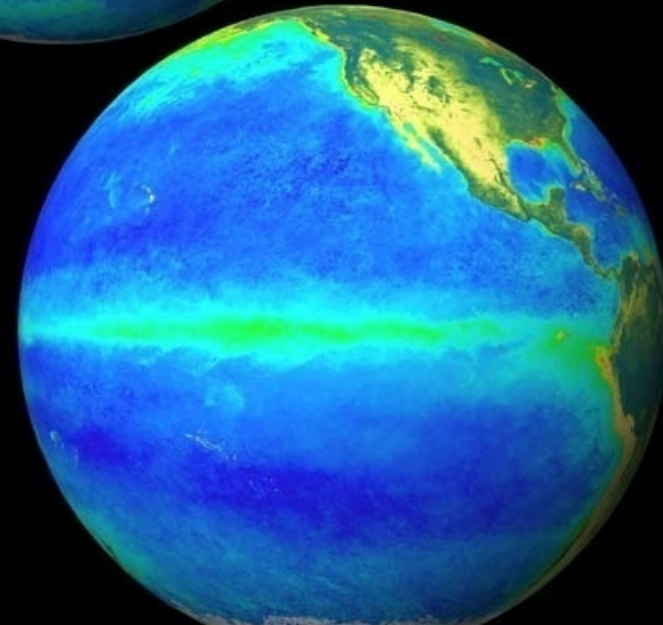
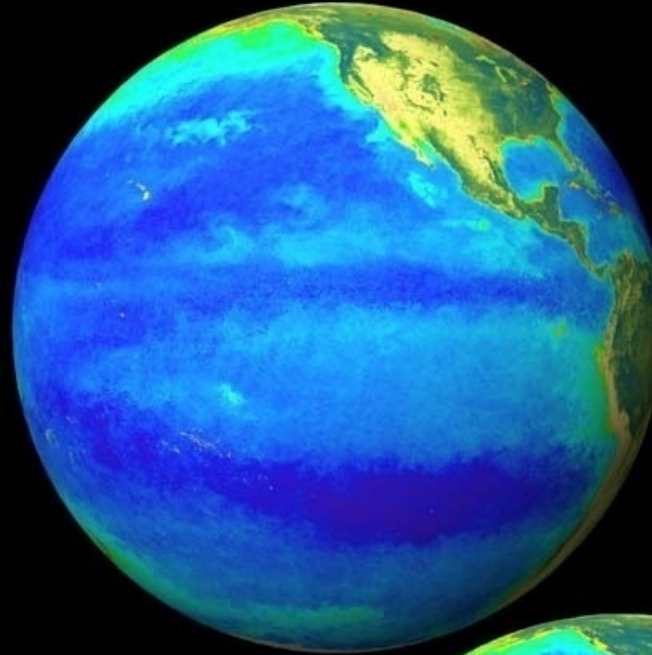
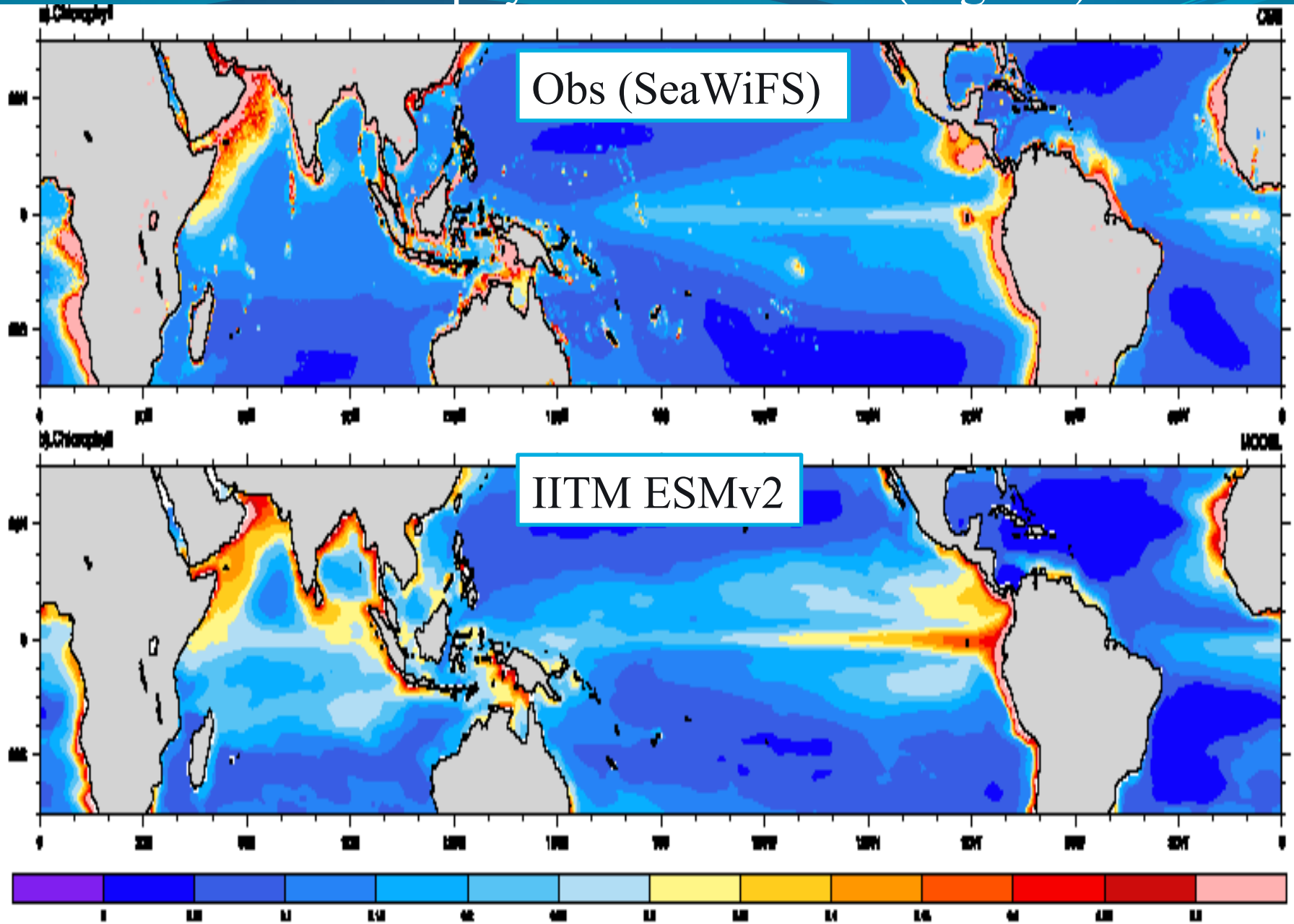


Image from SeaWiFS Project, NASA /

Chlorophyll Concentration (Mg m^{-3})

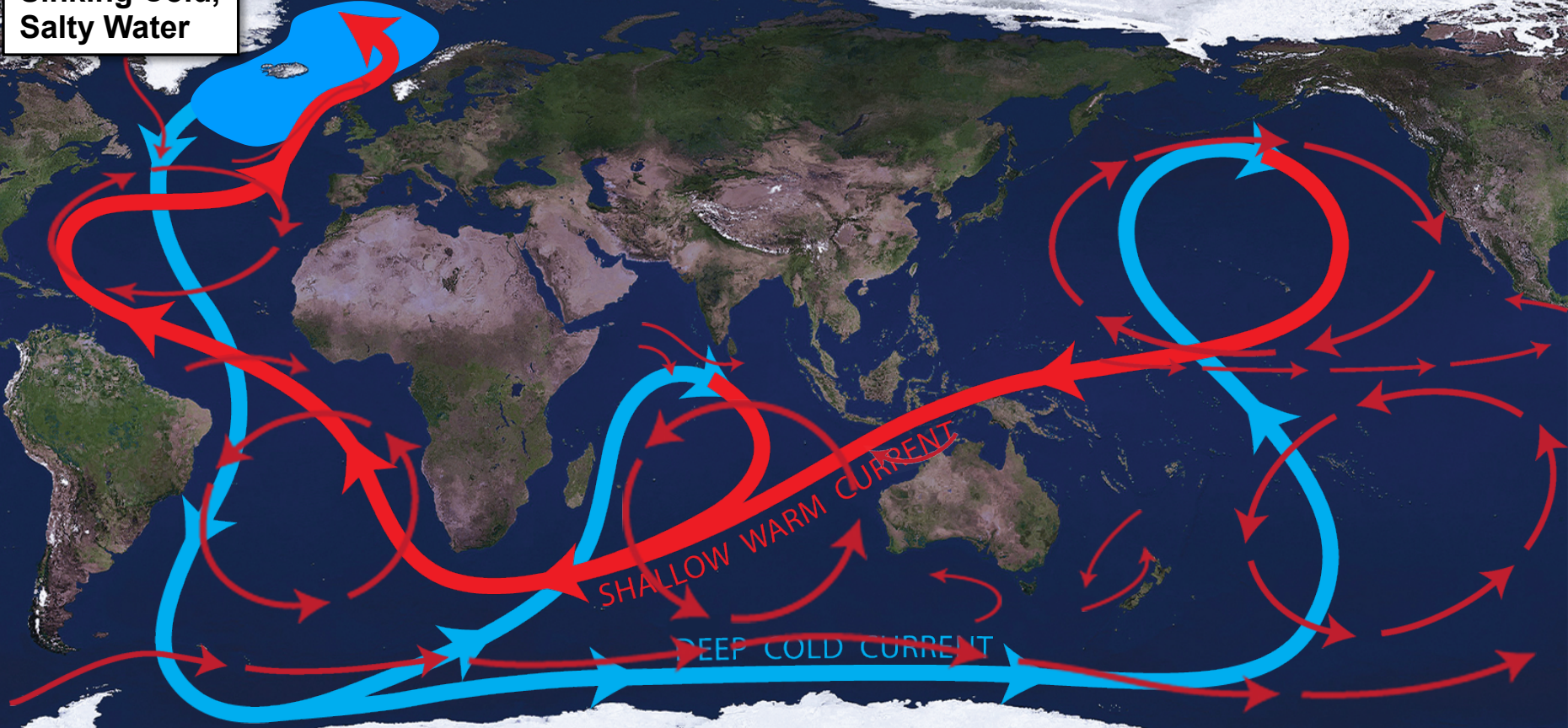


**Large-scale Ocean Circulation : Global
Overturning Circulation and Energy
Transport**

Thermohaline Circulation (THC)

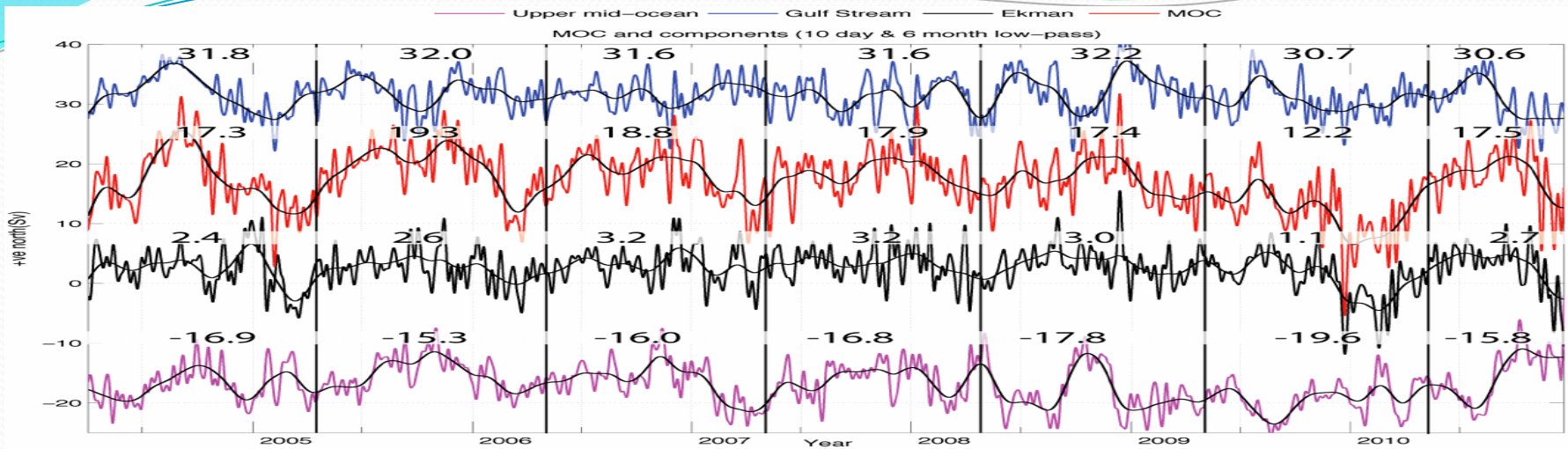
Global Conveyor Belt

Sinking Cold,
Salty Water



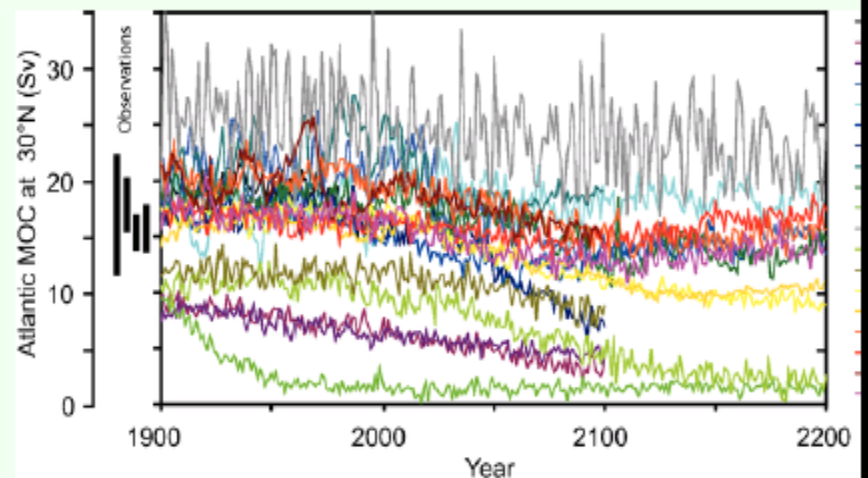
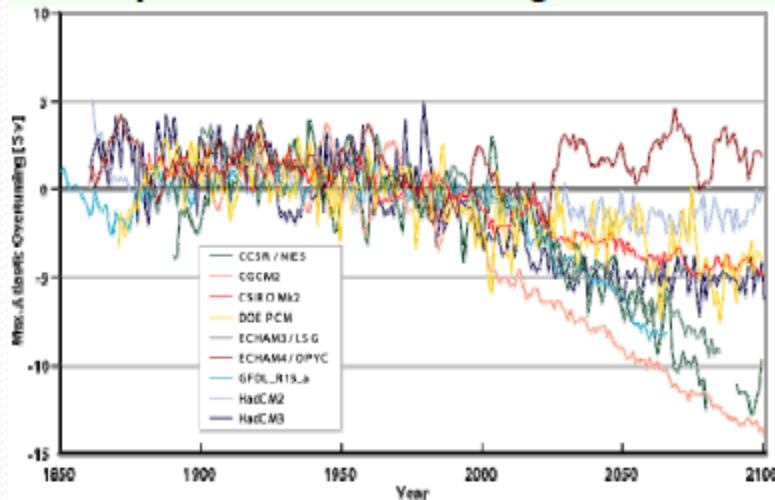
- . THC involves a deep overturning circulation driven by contrasts in density, and hence pressure, between different regions
- . THC is responsible for a large fraction of the $\sim O(1\text{PW})$ northward heat transport of the Atlantic Ocean

AMOC Variability in the Subtropical Atlantic (RAPID 26.5N)



- Natural (e.g. volcanic) and anthropogenic (e.g. GHG, aerosols) forcings can modify buoyancy fluxes, and influence AMOC

Response to GHG forcing: slow down due to warming and freshening Source: IPCC



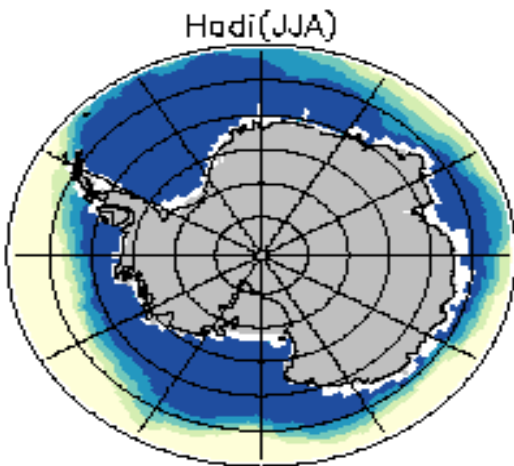
Sea-Ice concentration

Obs

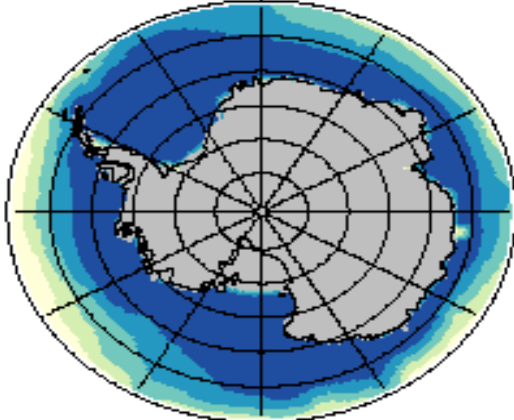
ESMv2

Improved simulation of
NH sea-ice during JJA

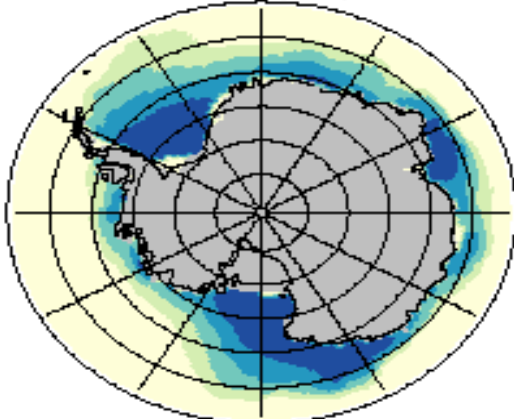
ESMv1



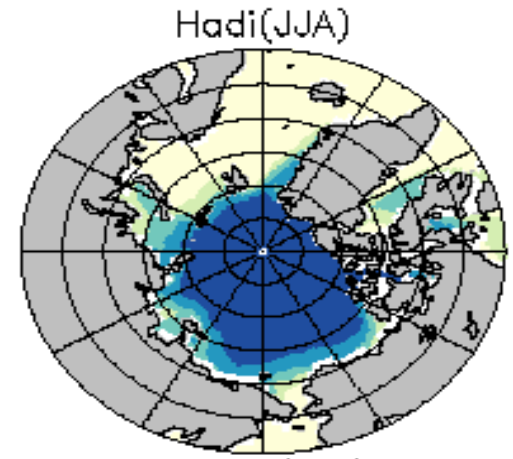
Hadim(JJA)



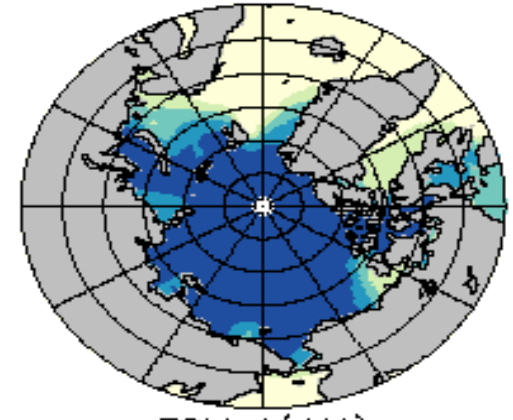
ESMv2(JJA)



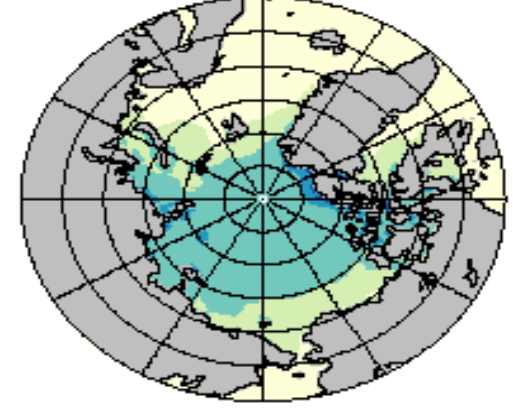
ESMv1(JJA)



Hadim(JJA)

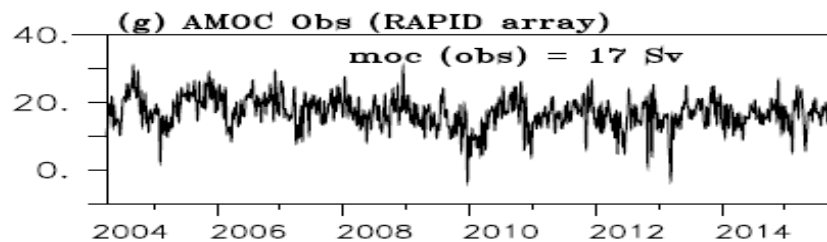
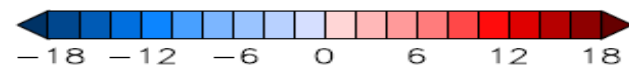
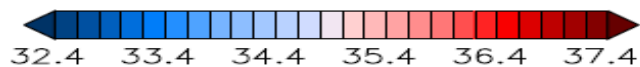
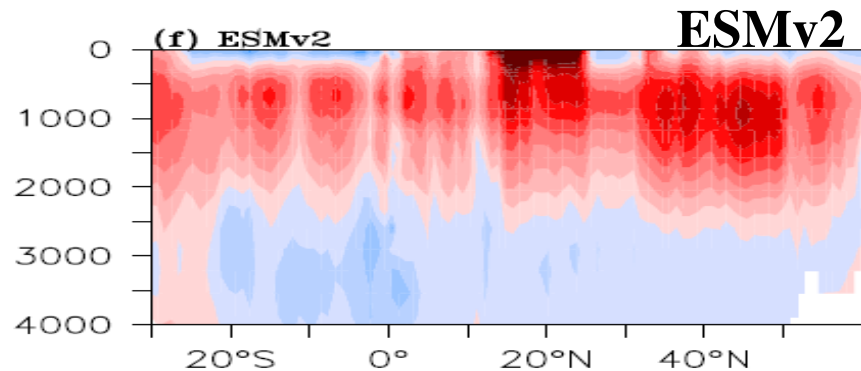
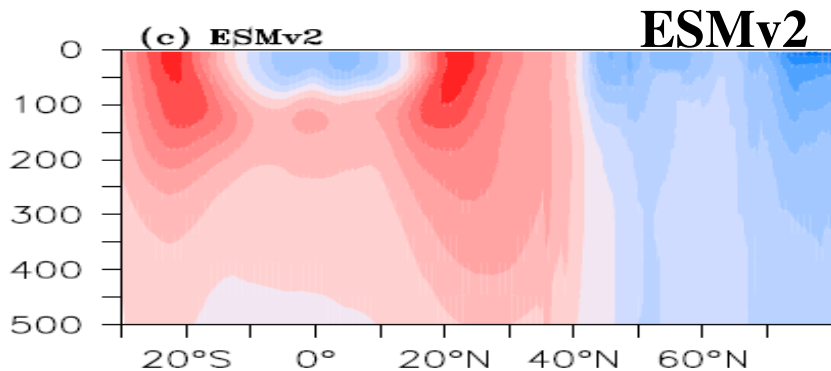
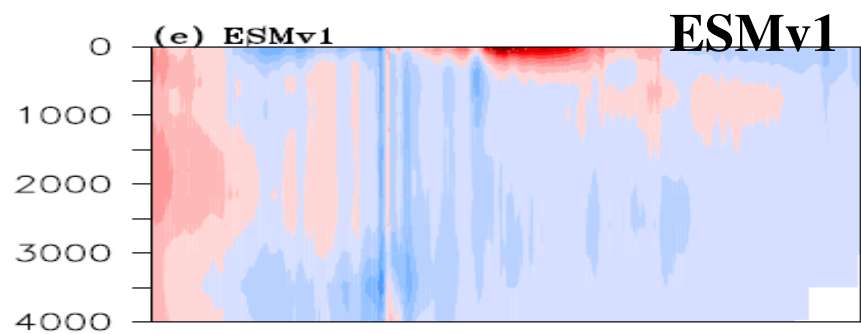
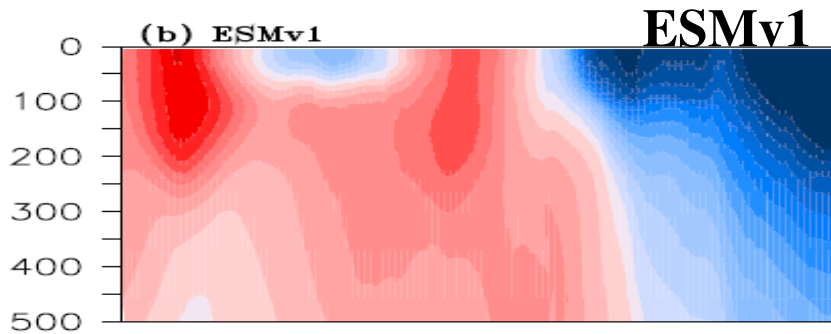
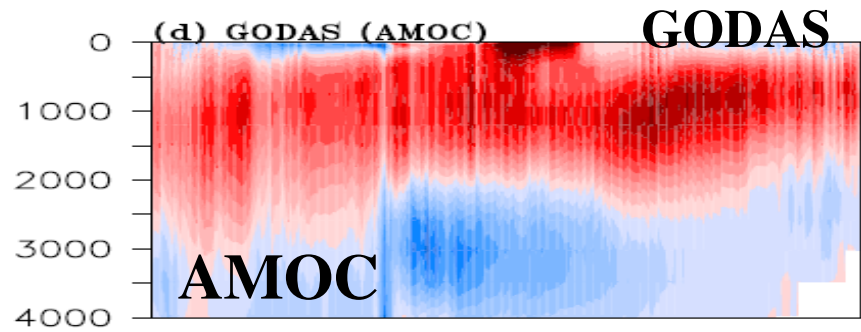
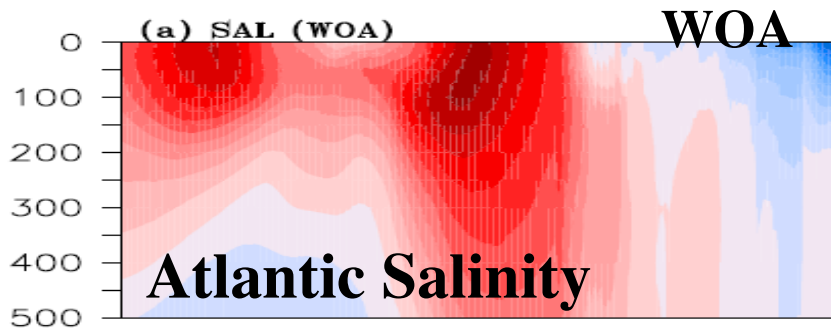


ESMv2(JJA)



ESMv1(JJA)





Annual Mean Meridional Heat Transport

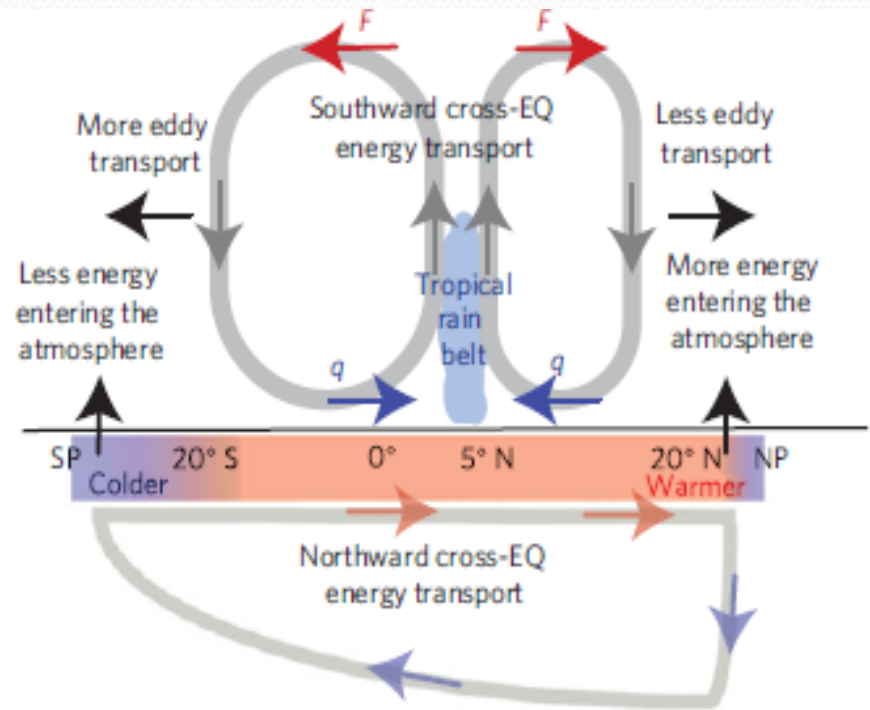
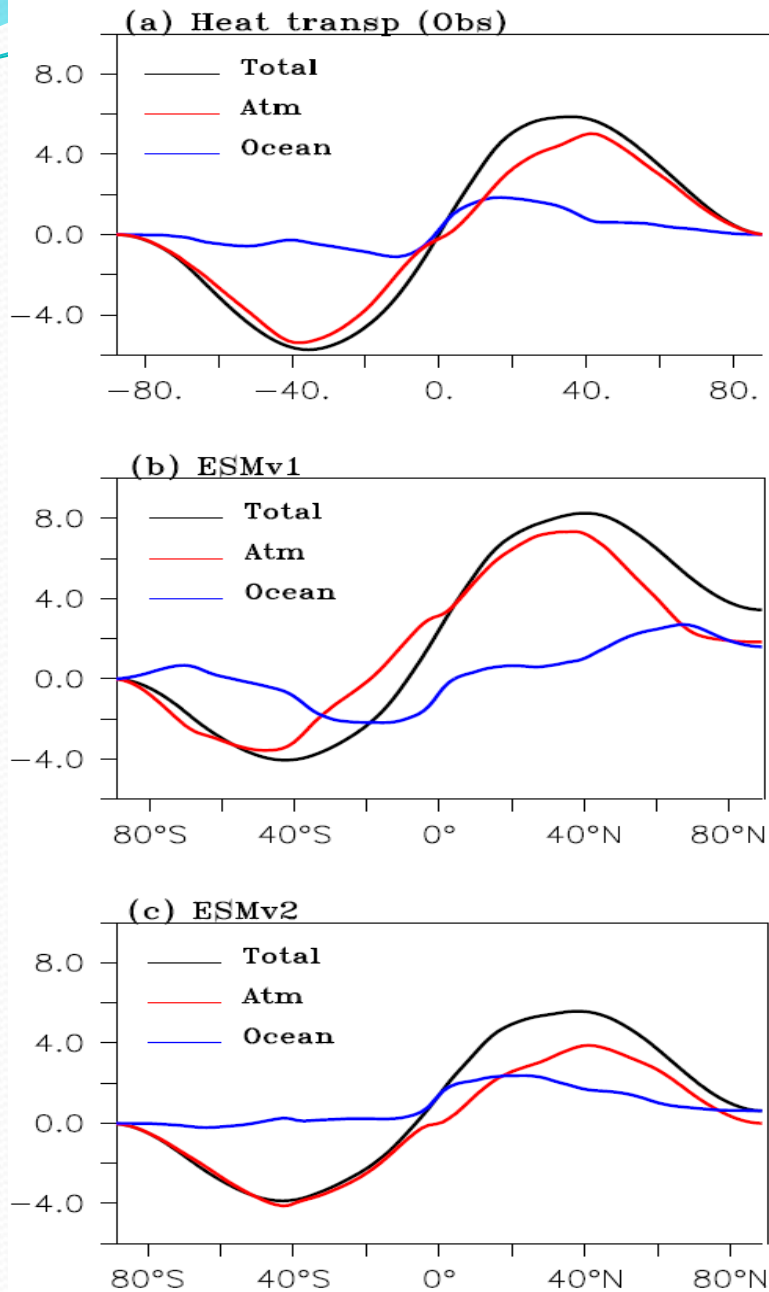
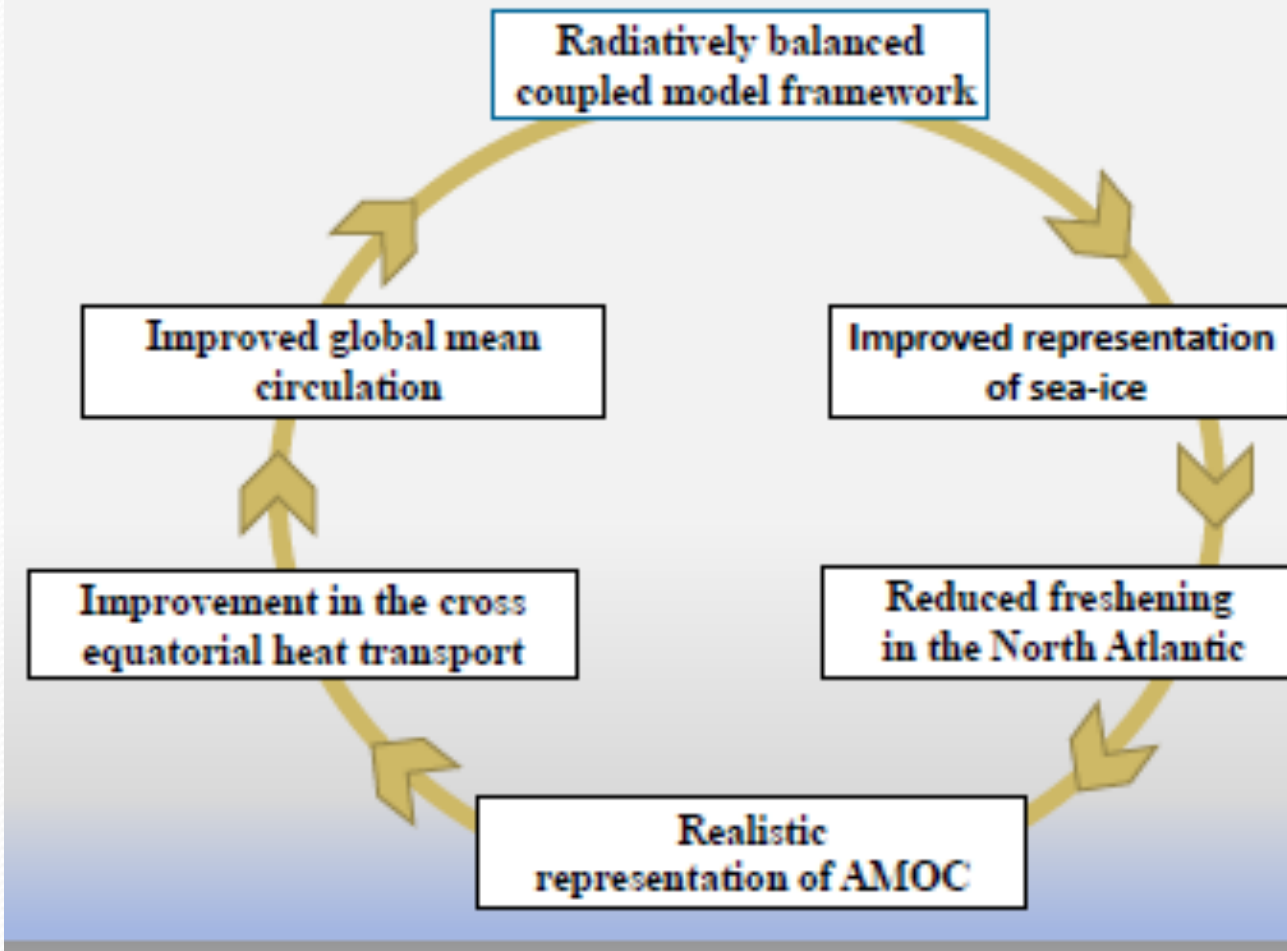


Figure 3 | Schematic of the role of the oceanic MOC in forcing the Northern Hemisphere maximum of tropical precipitation. Heat is released from the ocean to the atmosphere in the Northern Hemisphere owing to cross-equatorial OHT. The atmosphere responds through eddy energy transports in the extratropics, and a cross-equatorial Hadley circulation, which fluxes energy from the Northern Hemisphere to the Southern Hemisphere. The moisture transport by the Hadley circulation is in the opposite direction as the energy transport, so tropical precipitation moves northwards. SP, South Pole; NP, North Pole; cross-EQ, cross-equatorial; q , moisture transport; F , energy transport.

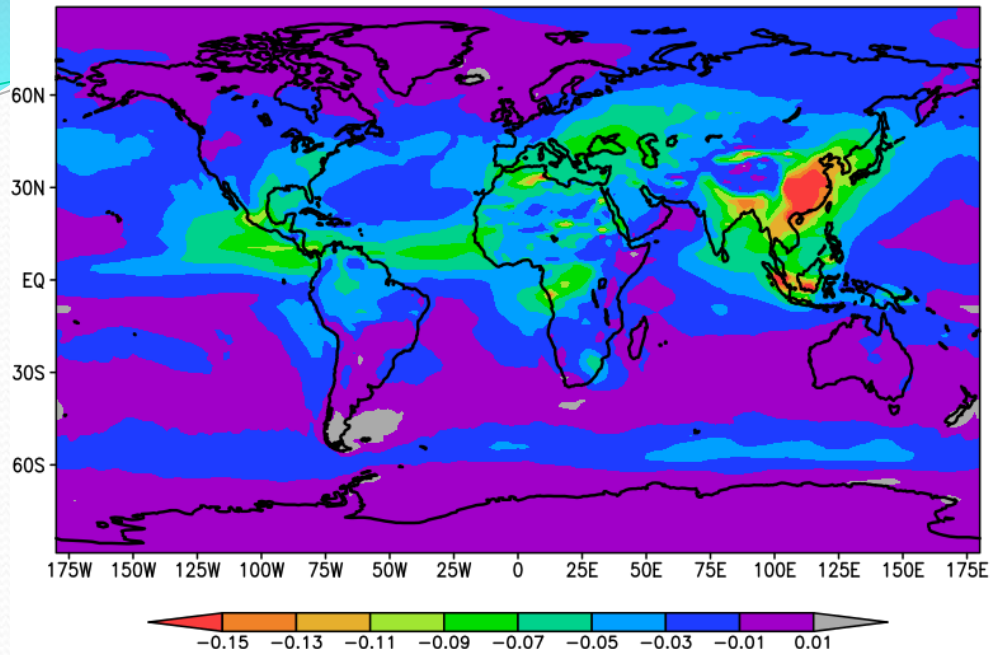
Frierson et al. 2013

Ocean-Atmosphere coupled feedbacks in IITM-ESM



Impact of aerosols and land-use land-cover changes on South Asian Monsoon

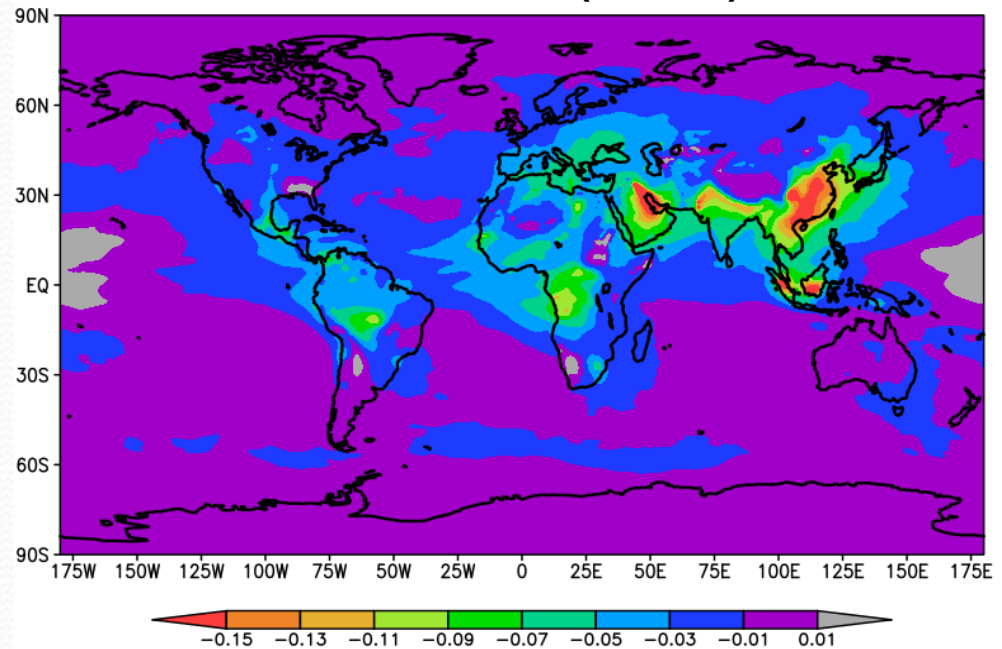
Calculated for IITM ESM



AOD difference map

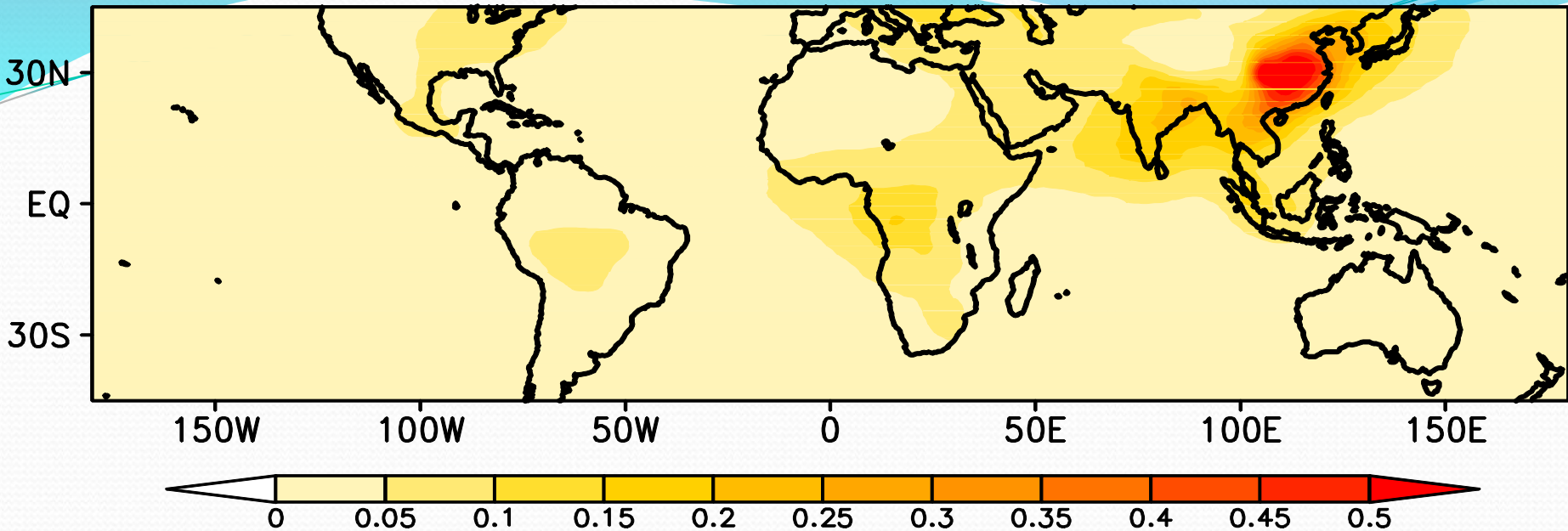
Preindustrial (1870)- Present (2005)

CESM (CAM5)

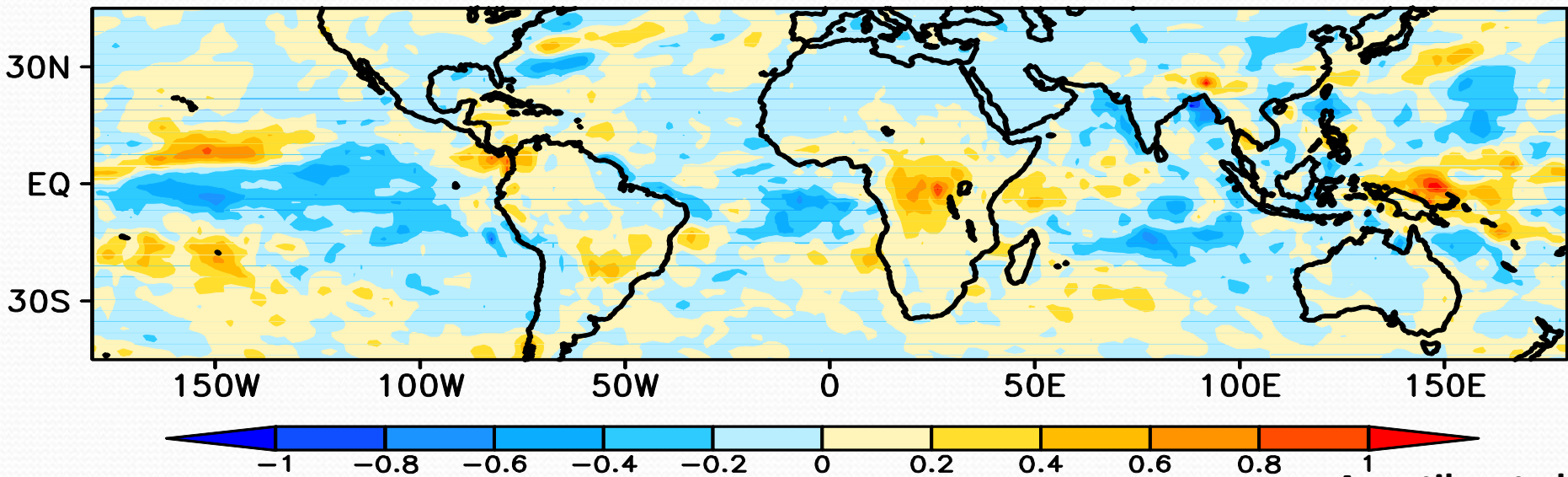


Courtesy: Ayantika, CCCR

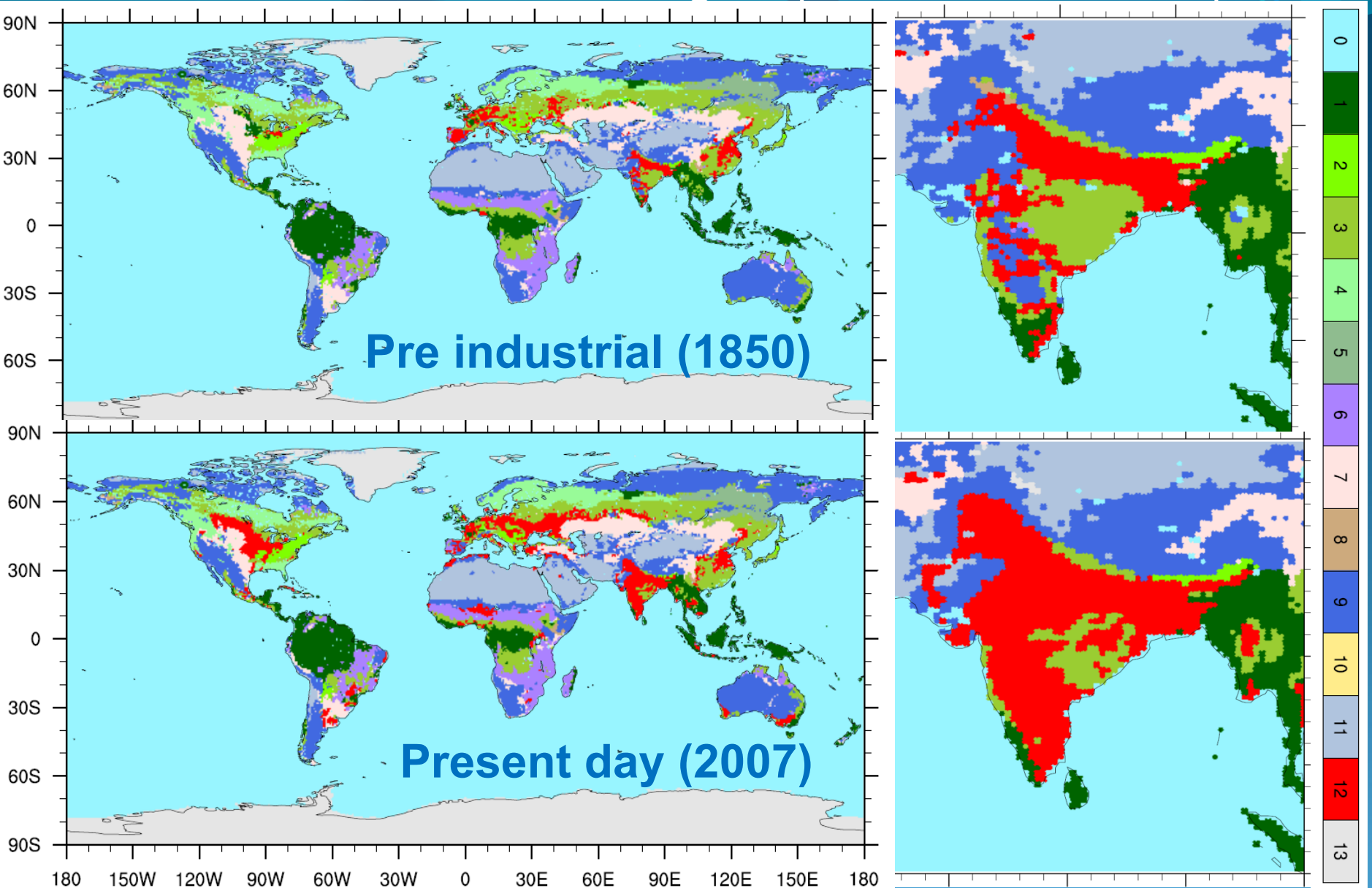
AOD-550nm(2005-1850)



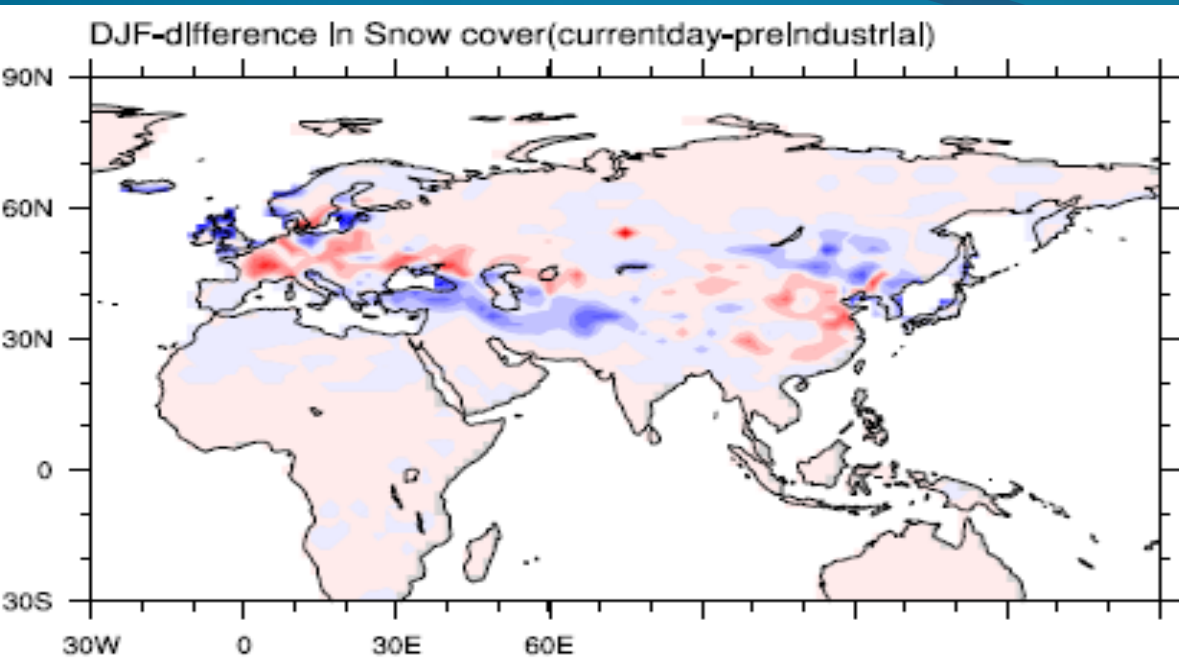
Precipitation(2005-1850)



Land use/land cover changes (Hurtt et al., 2015)

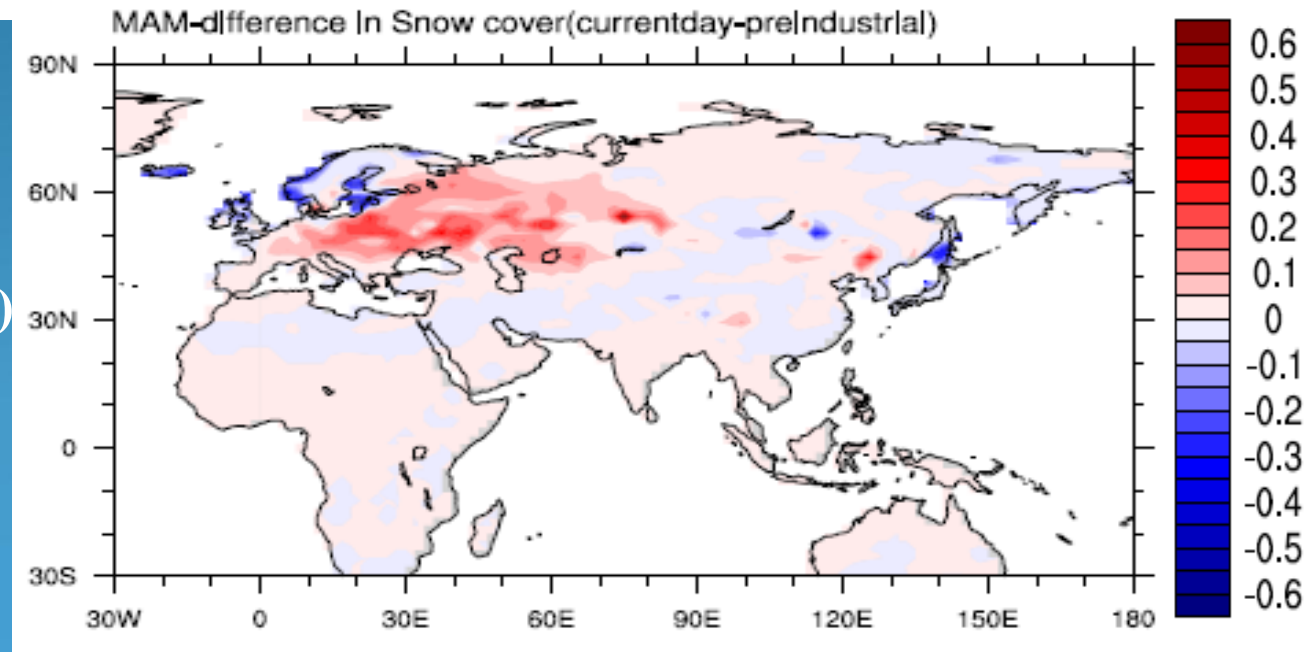


0 waterbodies	3 mixedforests	6 savannas	9 openshrubs	12 crops
1 Evergreen Broad	4 Evergreen Needle	7 grasslands	10 Tundra	13 snowice
2 DeciduousBroad	5 Deciduous Needle	8 shrubs	11 Barren	

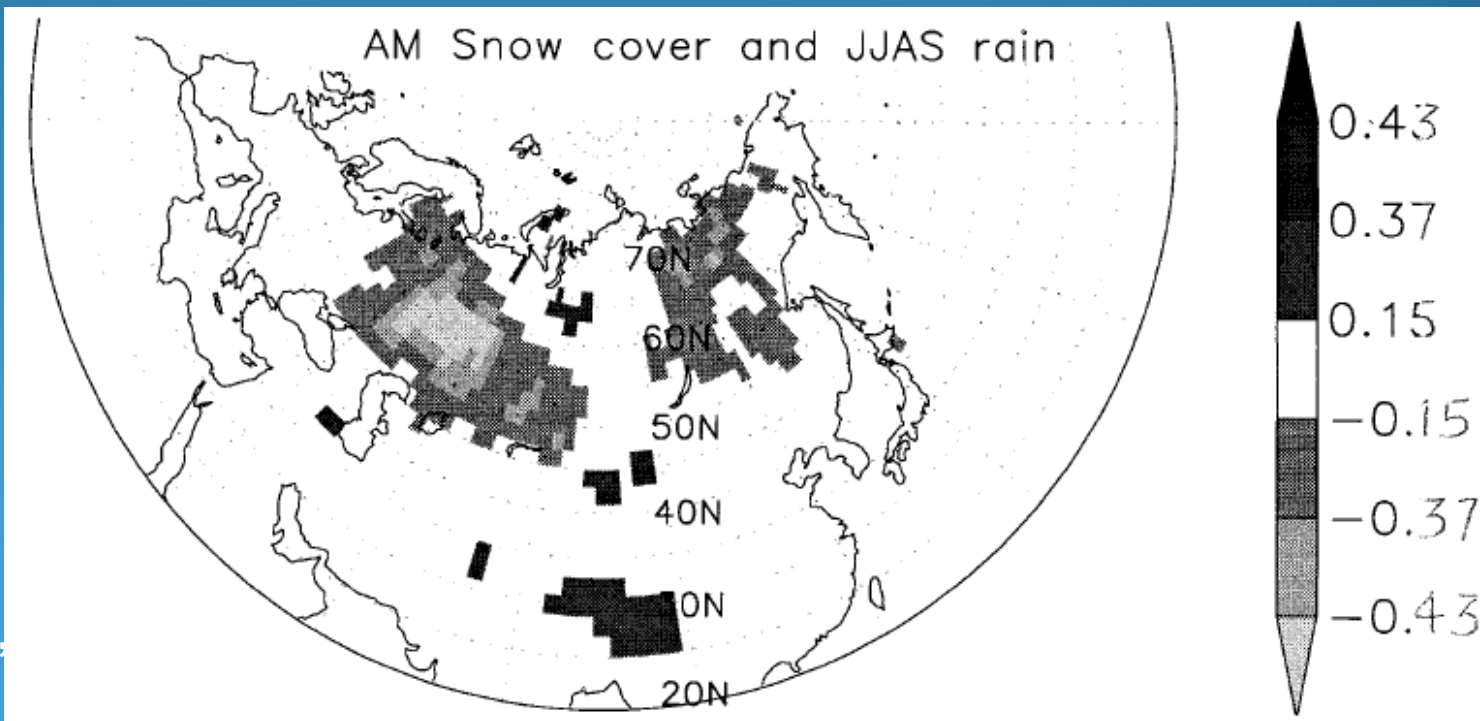
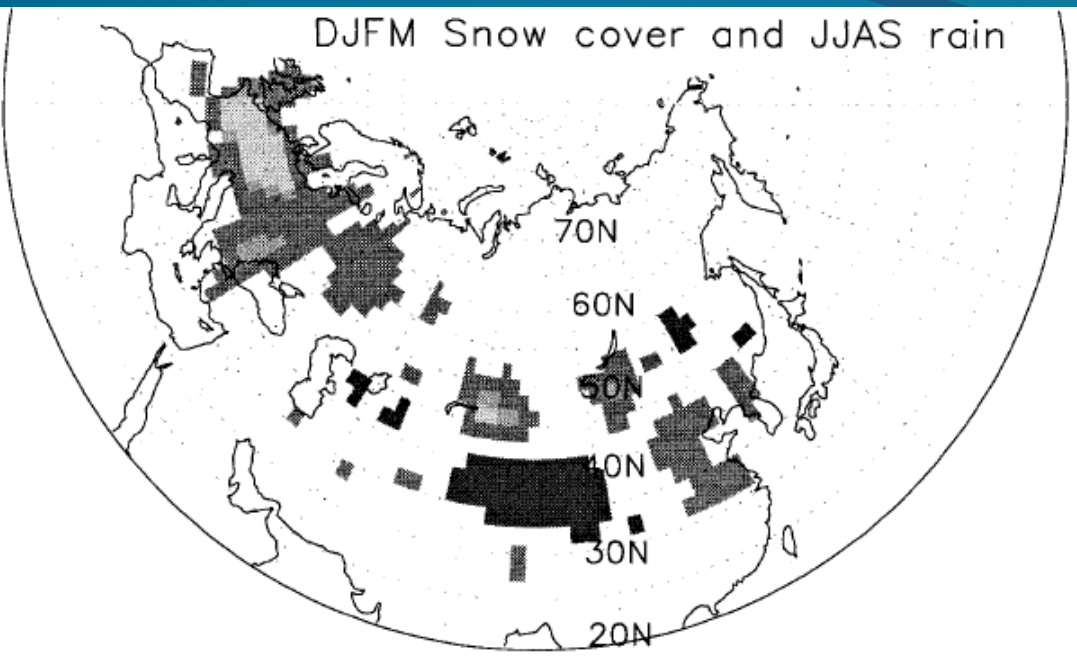


**Difference in Snow Cover
(Present – Pre-industrial)
Dec-Feb**

**(Present – Pre-industrial)
Mar-May**



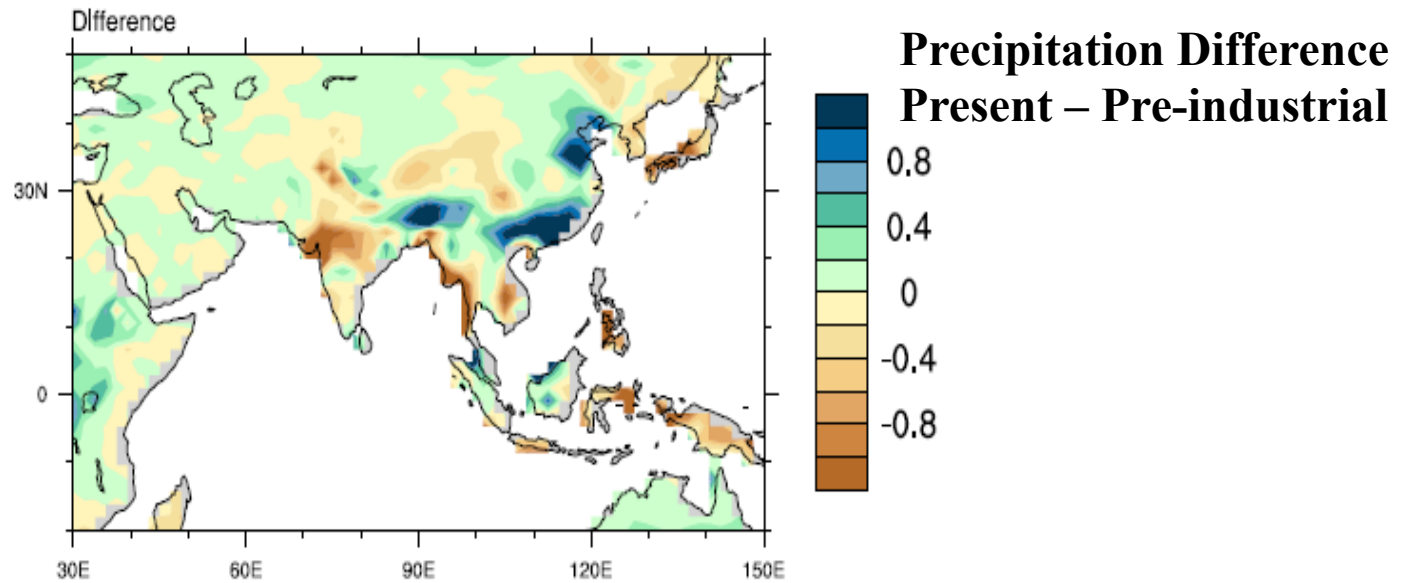
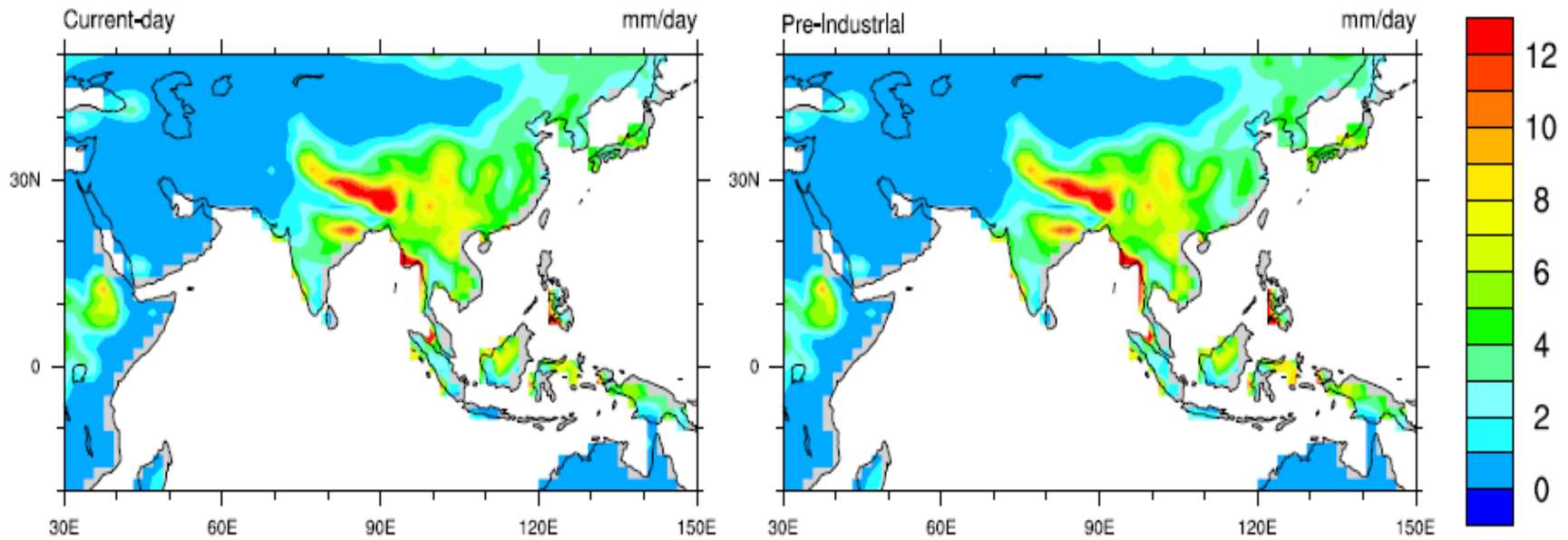
Eurasian snow cover and South Asian Monsoon



Bomzai & Shukla,
1998

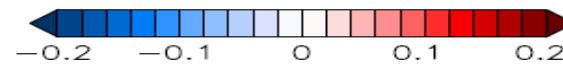
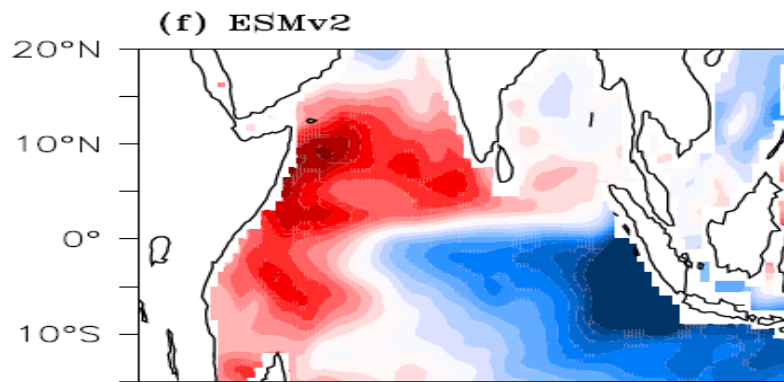
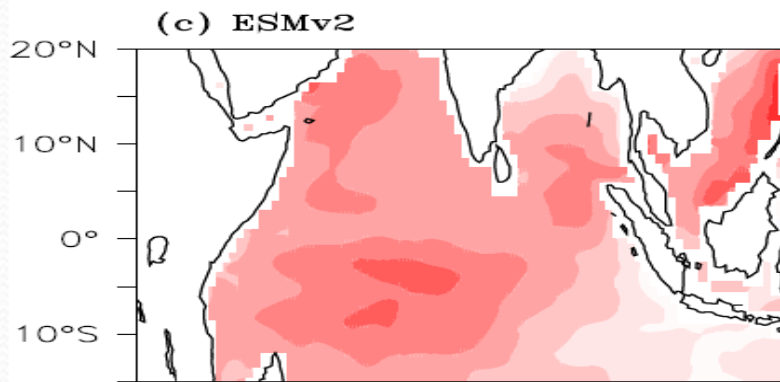
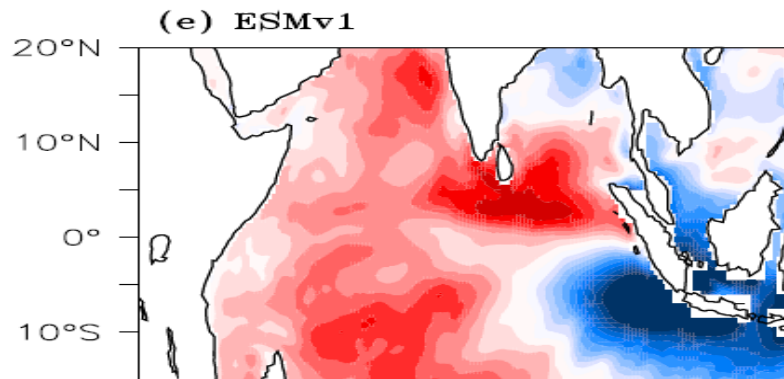
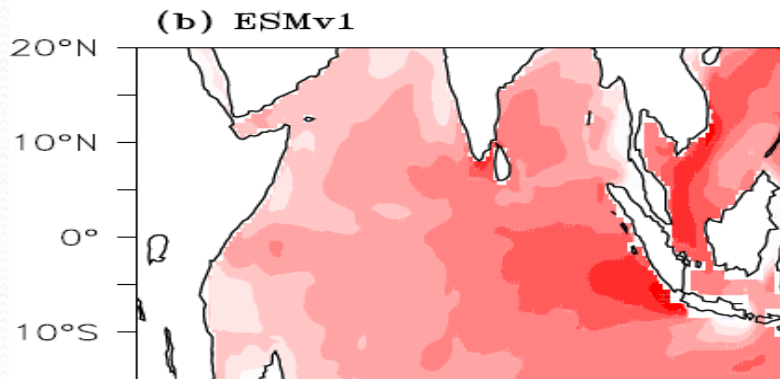
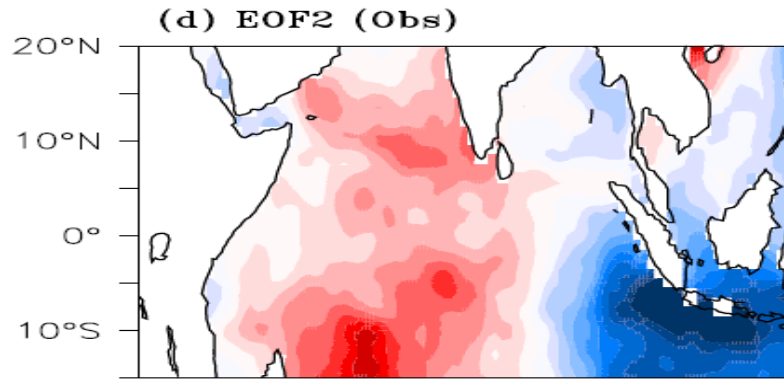
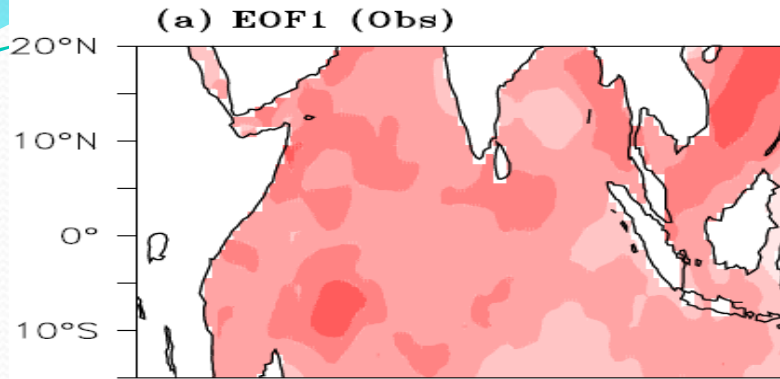
Precipitation (Present)

Precipitation (Pre-industrial)

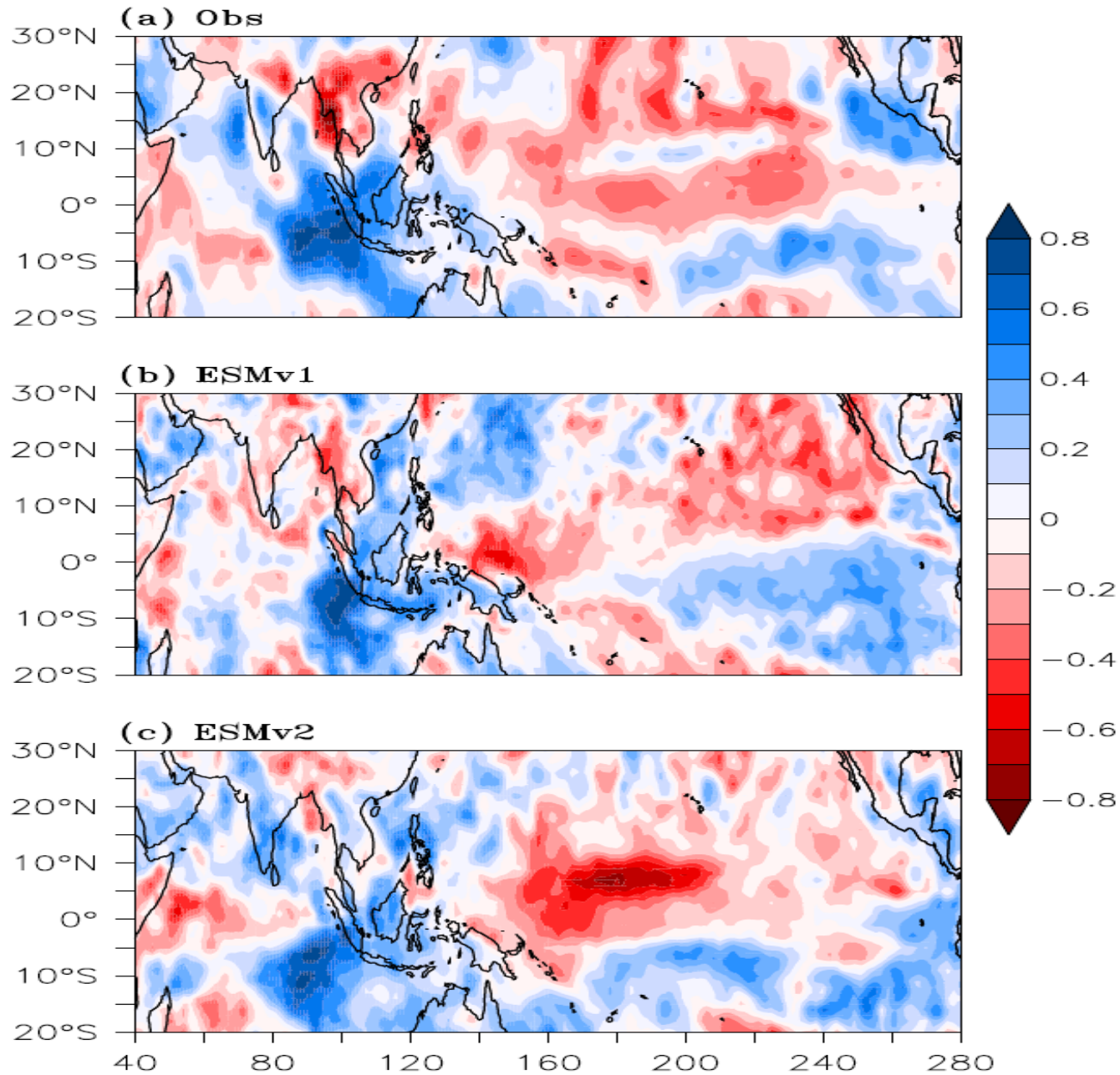


El Nino, IOD and its teleconnections to Indian Summer Monsoon Precipitation

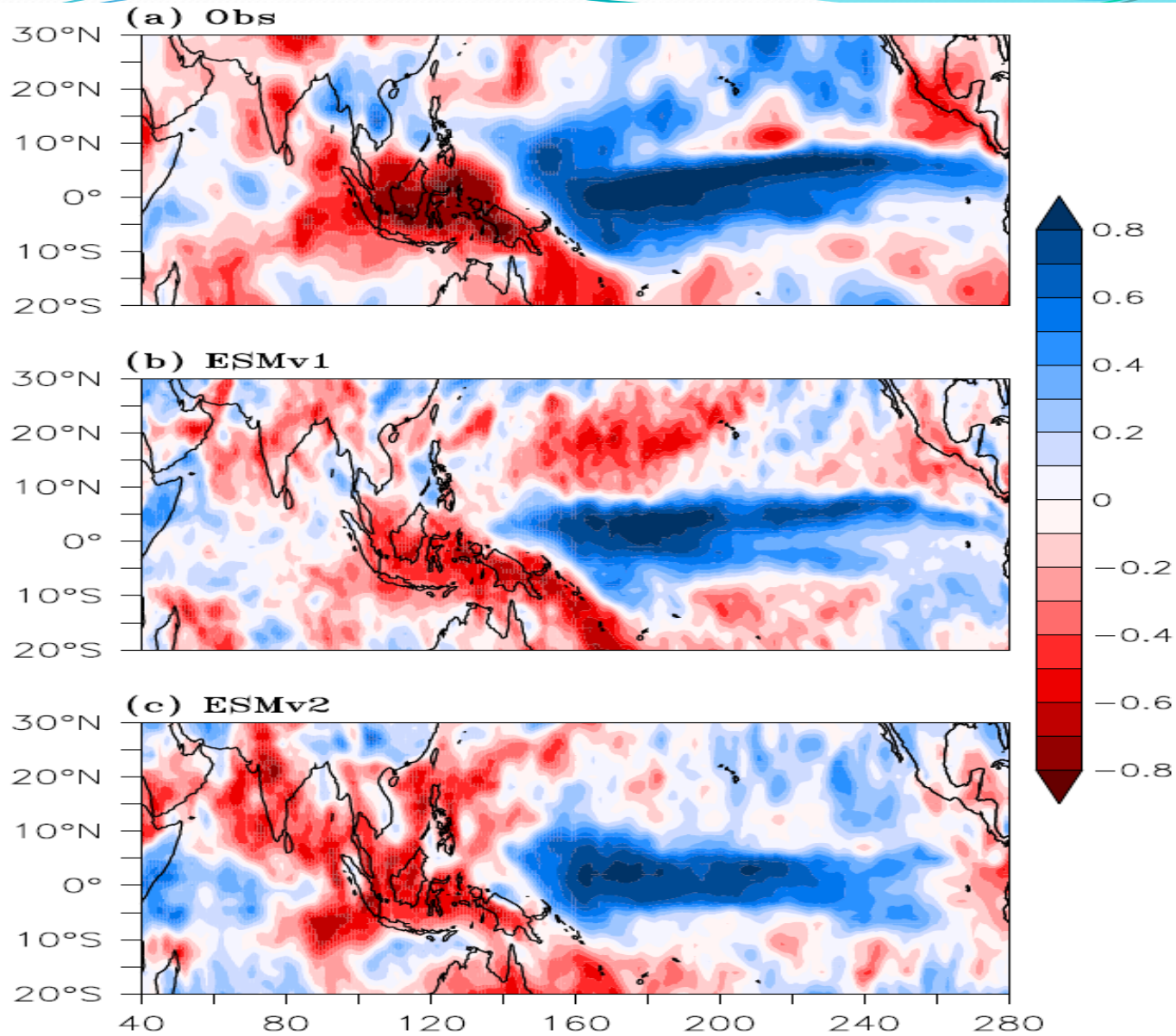
Tropical Indian Ocean Variability (IOBM & IOD)



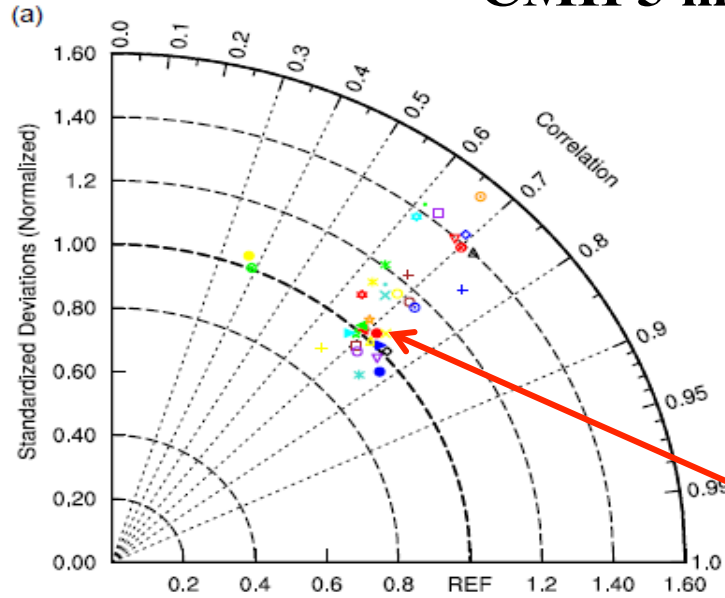
IOD and its teleconnection to South Asian Summer Monsoon Rainfall



ENSO and its teleconnection to South Asian Summer Monsoon Rainfall

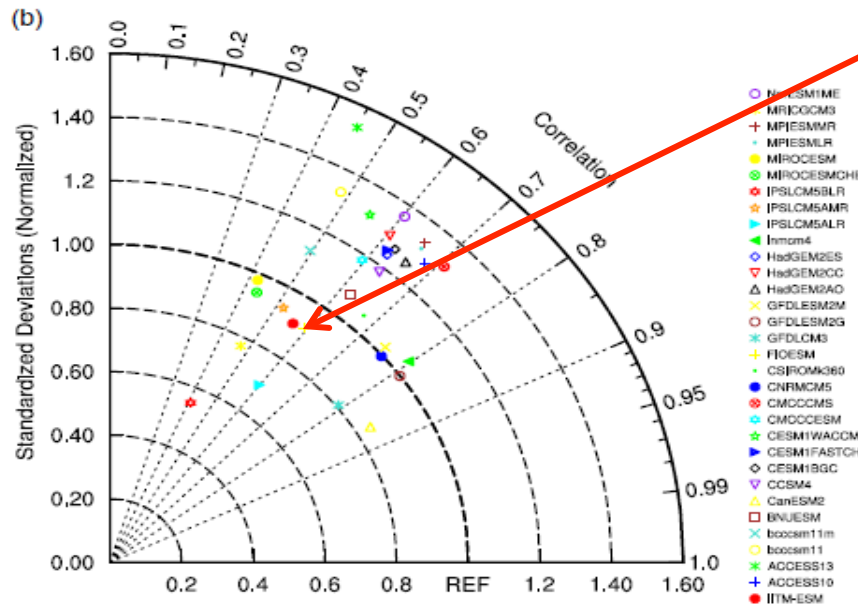


South Asian summer monsoon variability in IITM ESMv2 and CMIP5 models

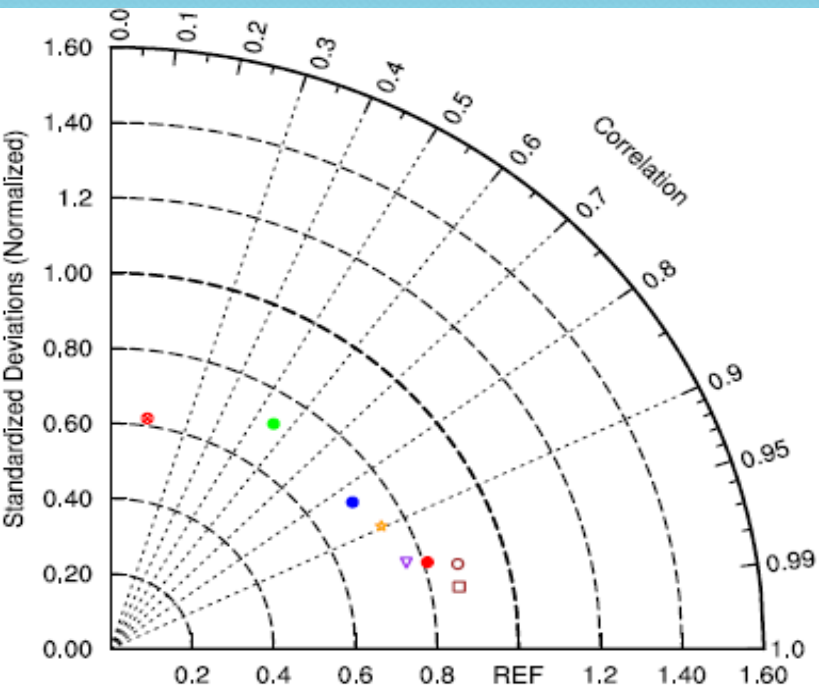


South and Southeast Asian Region
(65E-140E; 10S-40N)

IITM ESMv2

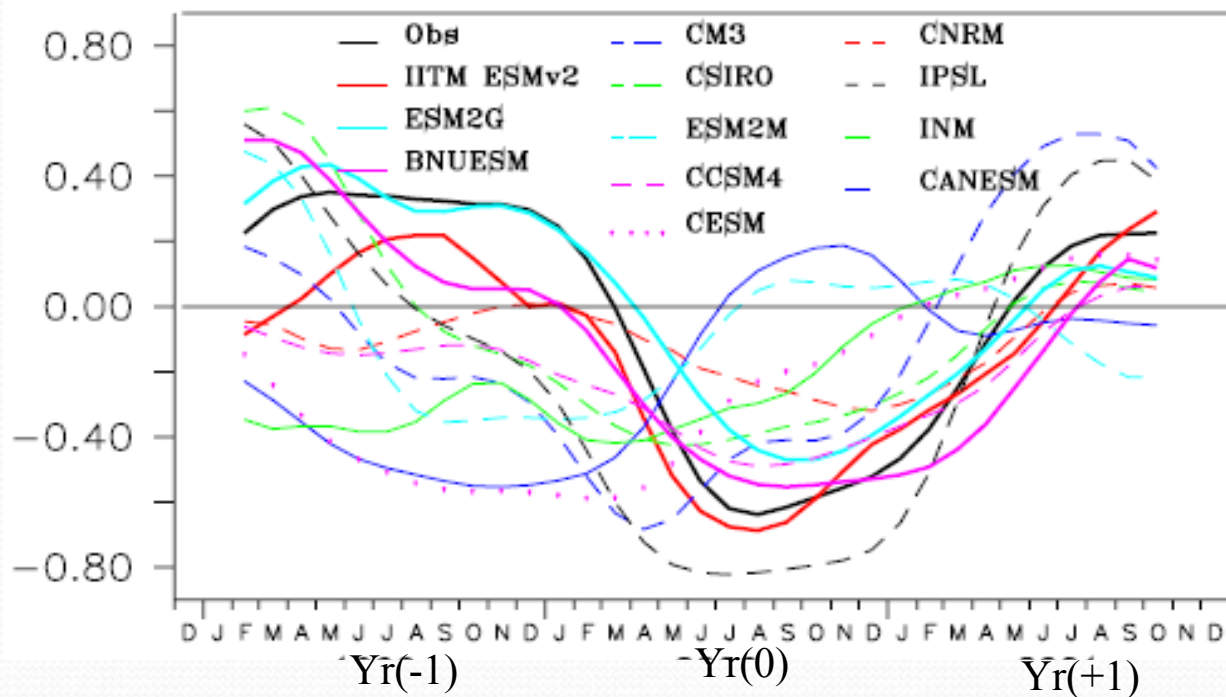


Indian land region



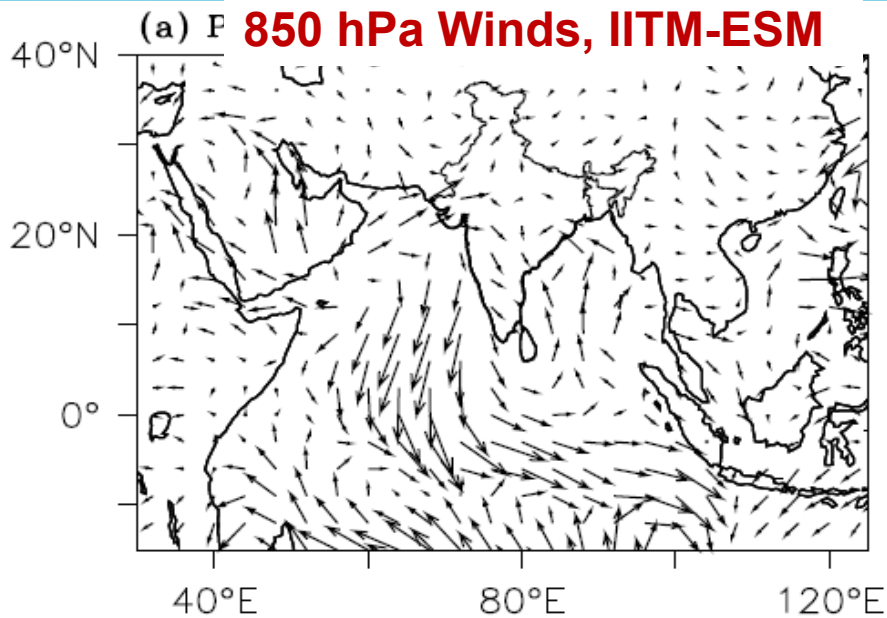
- NorESM1ME
- × MIROC5
- + MPIESM2R
- MPIESM2R
- MIROC5
- MIROC5MCHE
- IPSLCM5BLR
- IPSLCM5AMR
- IPSLCM5ALR
- Inmcm4
- HadGEM2ES
- HadGEM2CC
- HadGEM2AO
- GFDL2M2M
- GFDL2M2G
- GFDL2M3
- FIOESM
- CSIRO Mk3.6.0
- CNRMCM5
- CMCCCM3
- CMCCCM3
- CESM1WACCM
- CESM1FASTCH
- CESM1BGC
- CCSM4
- CanESM2
- BNUESM
- × bccsm11m
- bccsm11
- × ACCESS13
- + ACCESS10
- IITM-ESM

ENSO-Monsoon Teleconnection

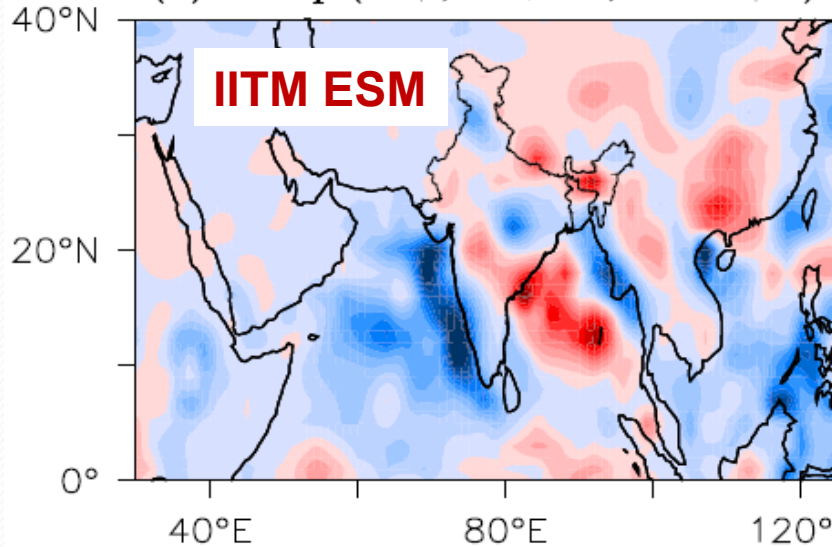


Present-Pre-industrial Simulation

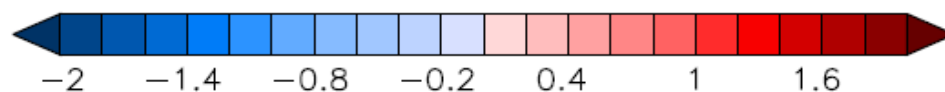
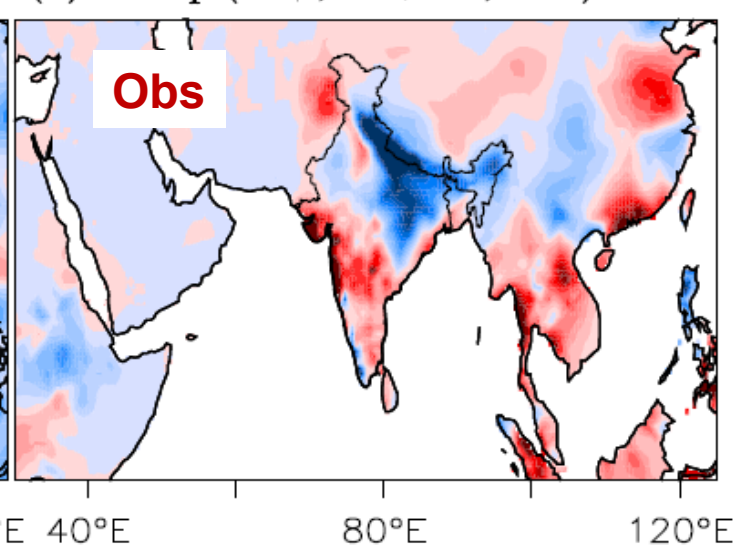
- Weakening Summer Monsoon Circulation
- Desiccation of Summer Monsoon Precipitation



(b) Precip (JJAS, Pres-PI, IITM-ESM)

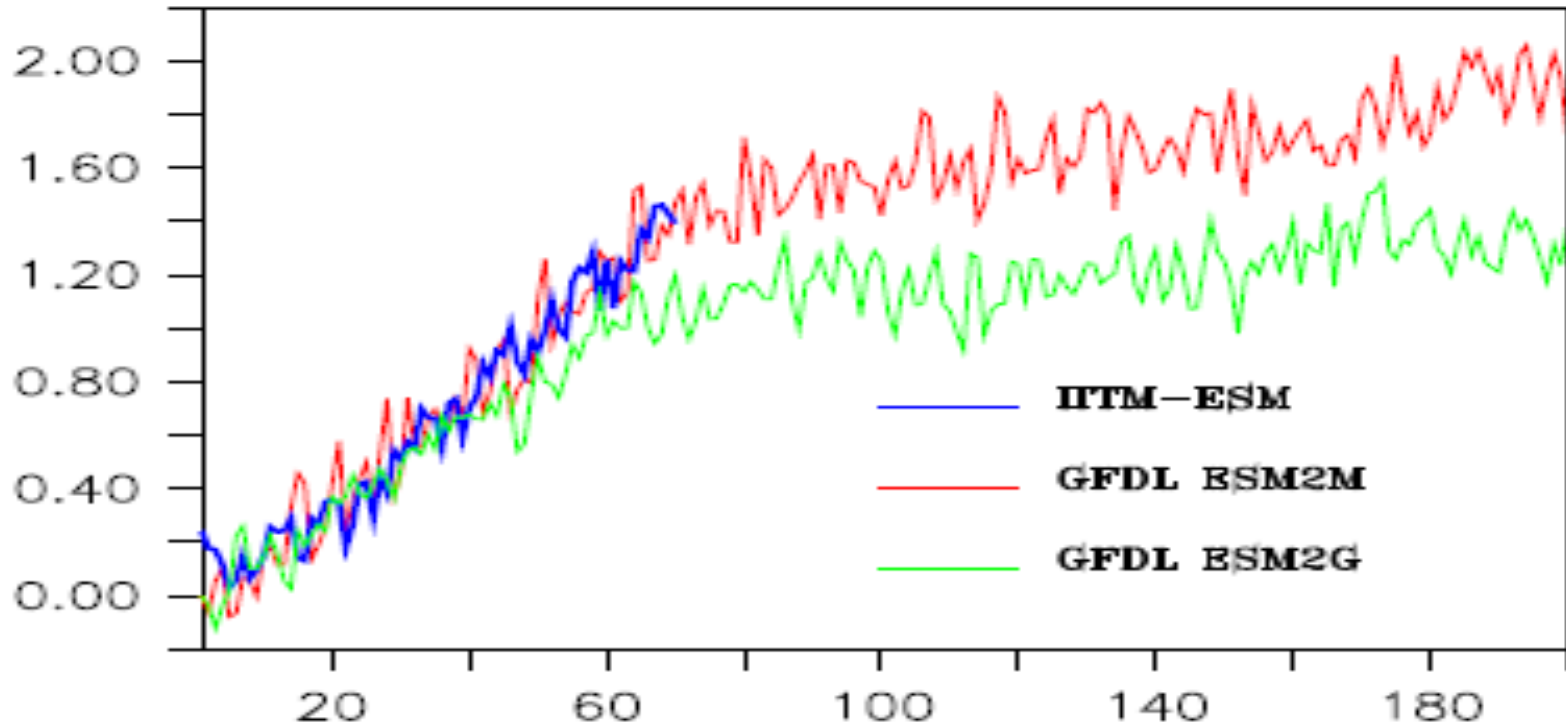


(c) Precip (JJAS, Pres-PI, CRU)



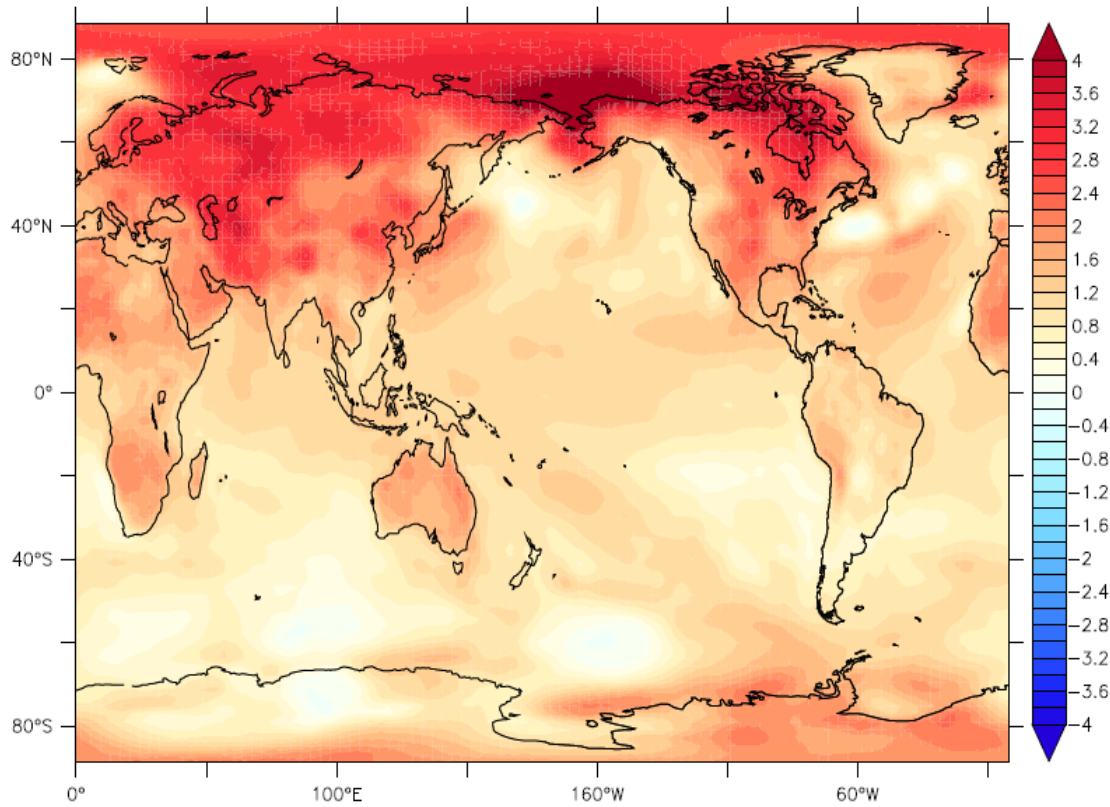
2xCO₂ response of IITM-ESM

TAS (2xCO₂)



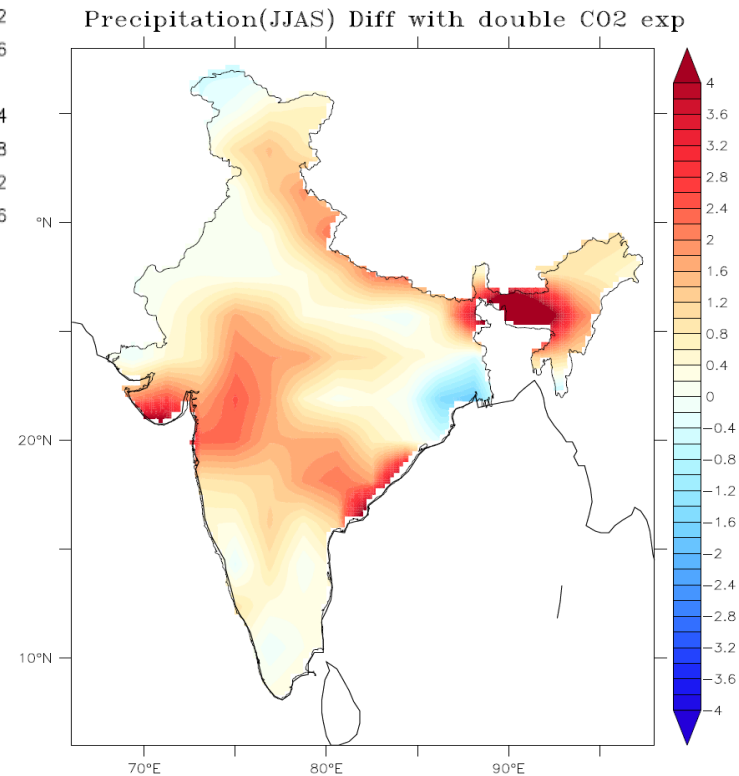
SST / SAT Anomaly: 2 x CO₂ experiment - PI Control

2xCO₂ response in IITM-ESM



Surface Temperature Anomaly

Precipitation Anomaly



Summary

The IITM-ESM has shown improvements in mean state & variability

- **Improved representation of mean monsoon features**
- **Land-atmosphere-ocean feedbacks**
- **ESNO & IOD, major drivers of Monsoon is well represented and teleconnections are robust**

These lead to improved mean and variability of summer monsoon precipitation over Asian region.

- **Need for improving the upper layer circulation features**



Thank You