

Future Changes in mean and extreme precipitation projected by 20-km and 60-km mesh MRI-AGCM ensemble simulations

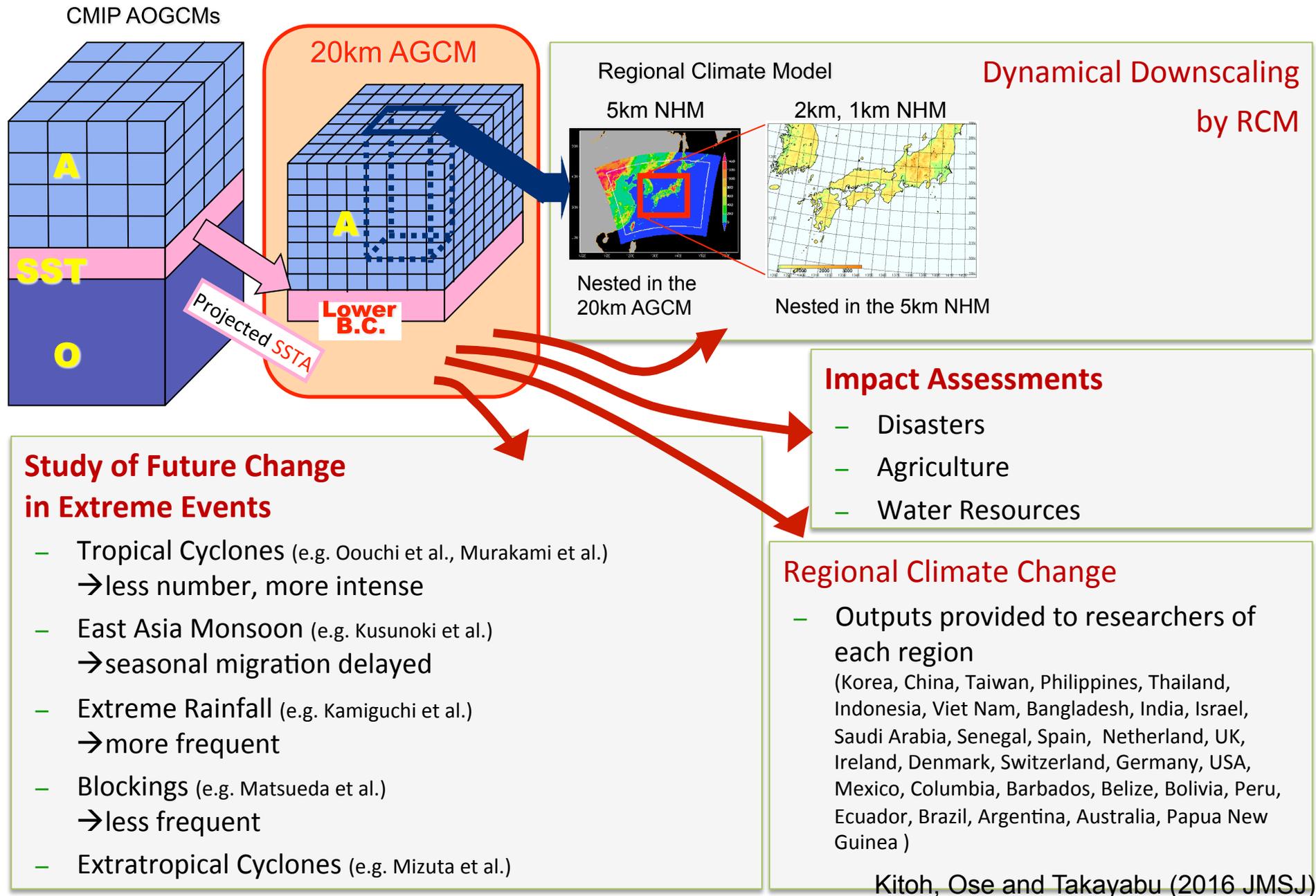
extreme precipitation and tropical cyclones (TCs)

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MRI-AGCM (60km, 20km) + NHRCM (5km, 2km)



History of MRI-AGCM3

MRI-AGCM3.0 (before 2007) (Mizuta et al. 2006; Oouchi et al. 2006)
developed from JMA operational NWP model
first 20km climate model which simulates for multi-decades



very minor change

MRI-AGCM3.1 (since 2007) (Kitoh et al. 2009; Murakami et al. 2011)
AMIP-type experiment



introducing new parameterization schemes etc.

MRI-AGCM3.2 (current model) (Mizuta et al. 2012)



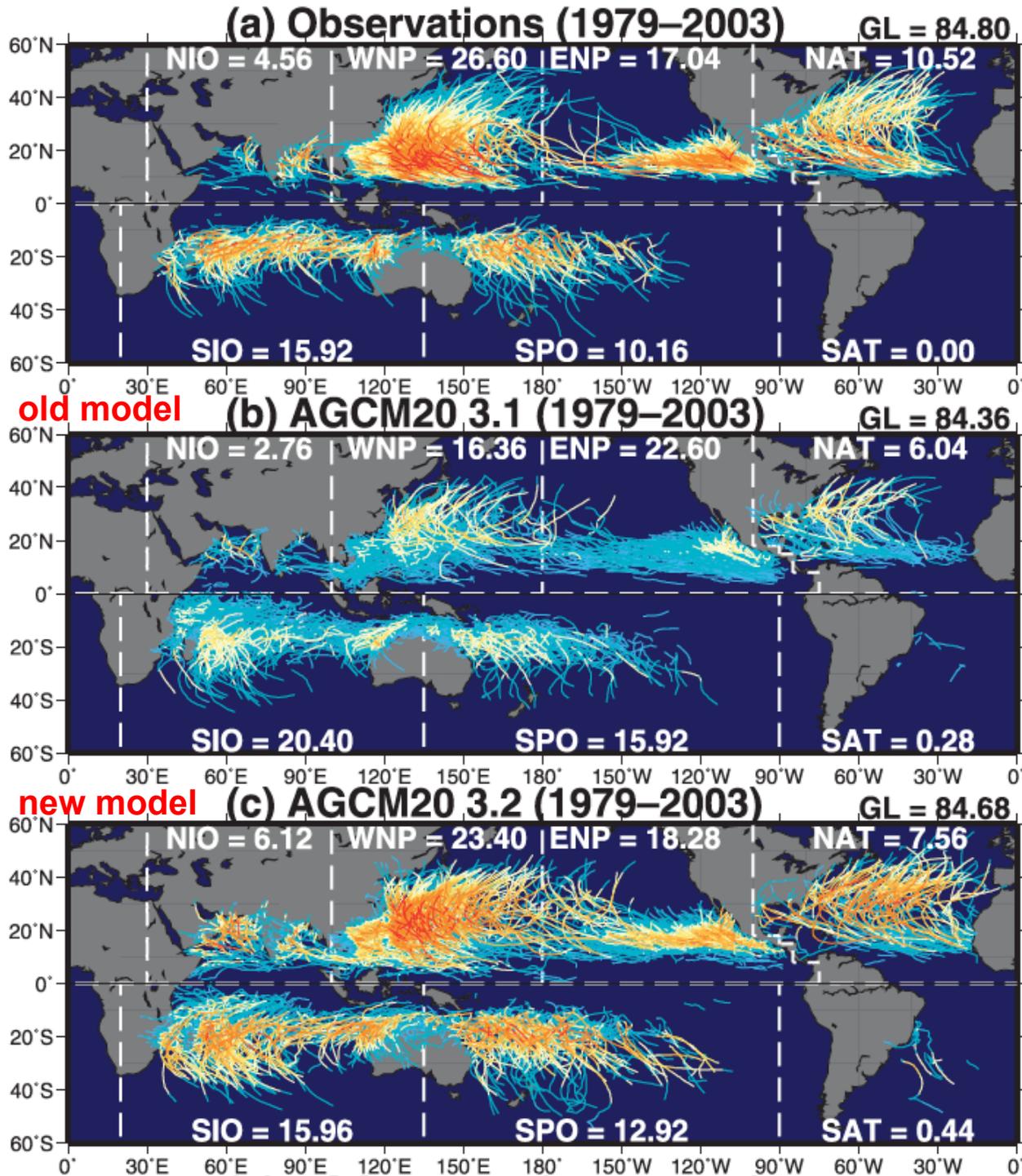
Reductions of the model biases of:

- insufficient precipitation amounts over the W. Pacific
- geographical distribution of tropical cyclones
- overestimated weak rain, underestimated heavy rain
- resolution dependence in terms of global-scale climate



Ensemble simulations with lower-resolution version

Tropical cyclone (TC)



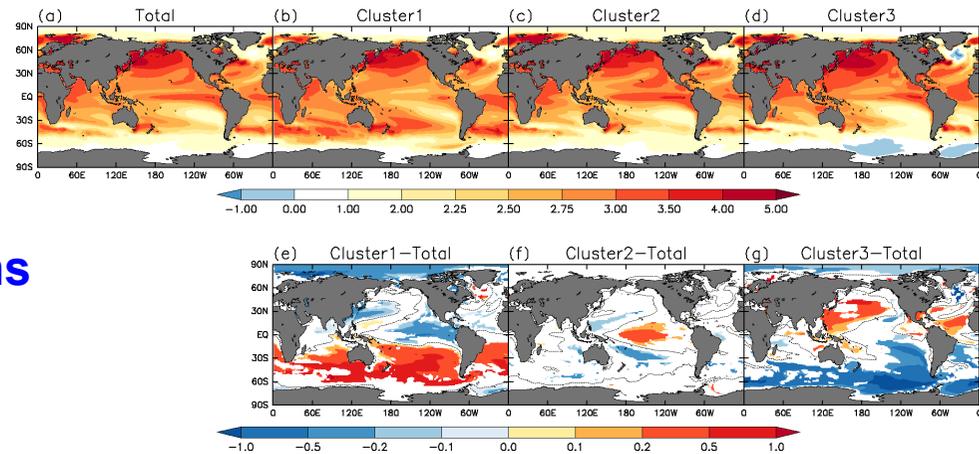
- Simulated TC number in the WNP is underestimated
- TC intensity is weak compared with observations

Model development

improved

MRI-AGCM3.2 experiment design

- **20km-mesh AGCM: 4 projections**
 - 4 different Δ SST patterns
- **60km mesh AGCMs: 12 projections**
 - 4 different Δ SST patterns
 - 3 different cumulus schemes



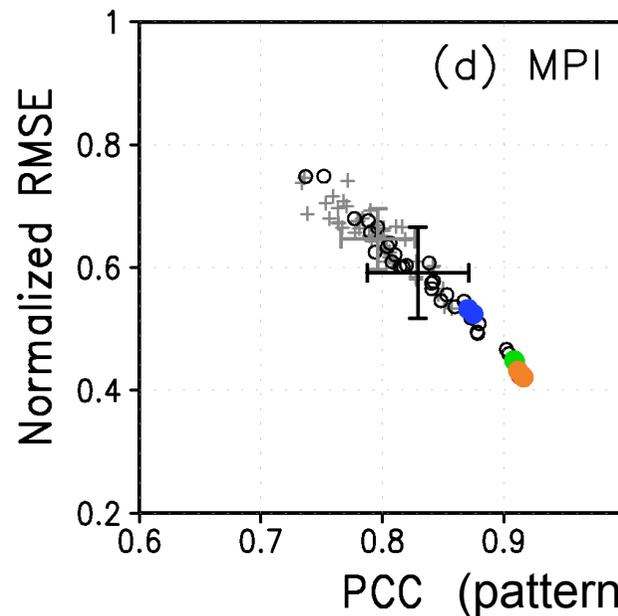
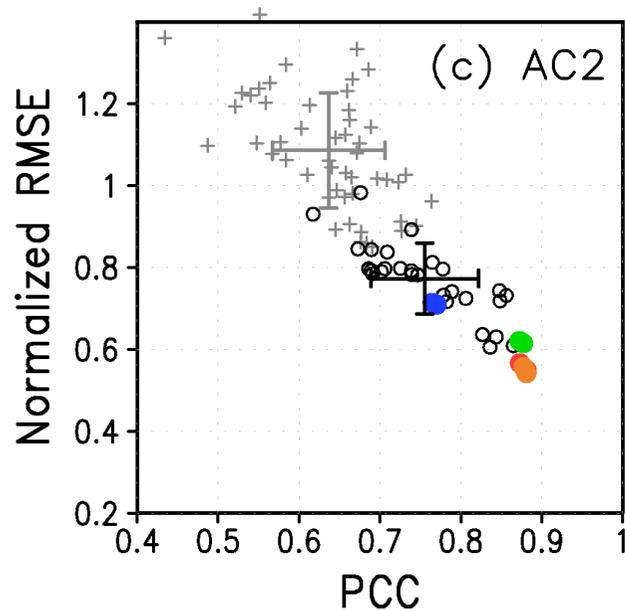
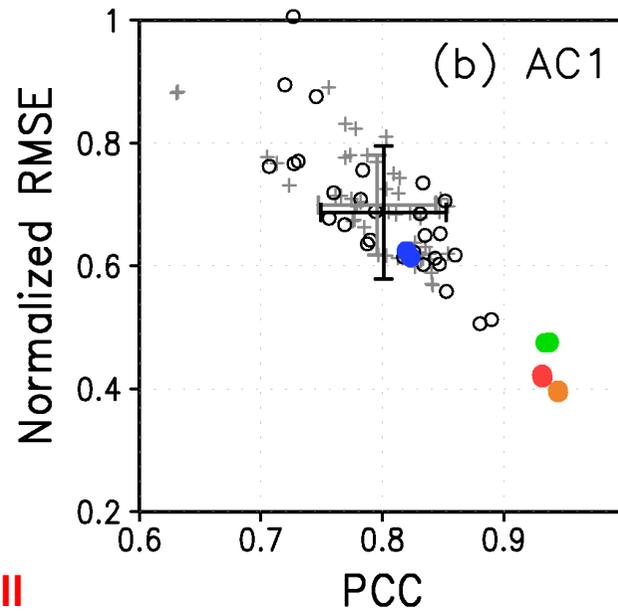
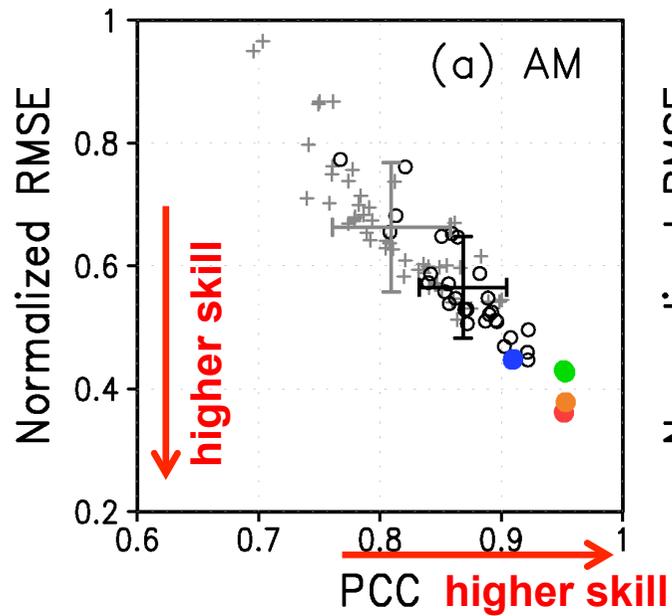
Δ SST: based on CMIP5-AOGCM projections (RCP8.5 scenario) (Mizuta et al. 2014)

Duration	SST	Cumulus convection scheme			
		Yoshimura (YS)	Arakawa-Schubert (AS)	Kain-Fritsch (KF)	
Present 1984-2003	Observation	SPA	HPA	HPA_as	HPA_kf
		SPA_m01	HPA_m02	HPA_as_m02	HPA_kf_m02
Future 2080-2099	All	SFA_rcp85	HFA_rcp85	HFA_as_rcp85	HFA_kf_rcp85
	Cluster 1	SFA_rcp85_c1	HFA_rcp85_c1	HFA_as_rcp85_c1	HFA_kf_rcp85_c1
	Cluster 2	SFA_rcp85_c2	HFA_rcp85_c2	HFA_as_rcp85_c2	HFA_kf_rcp85_c2
	Cluster 3	SFA_rcp85_c3	HFA_rcp85_c3	HFA_as_rcp85_c3	HFA_kf_rcp85_c3

20km

60km

Performance of global-scale monsoon precipitation



Monsoon metrics

(B. Wang et al. 2011)

AM: annual mean

AC1: summer - winter

AC2: spring - autumn

MPI: seasonal difference
divided by ann. mean

Area: 45S-45N

Ref.: (GPCP+CMAP)/2

● 20kmYS

● 60kmYS

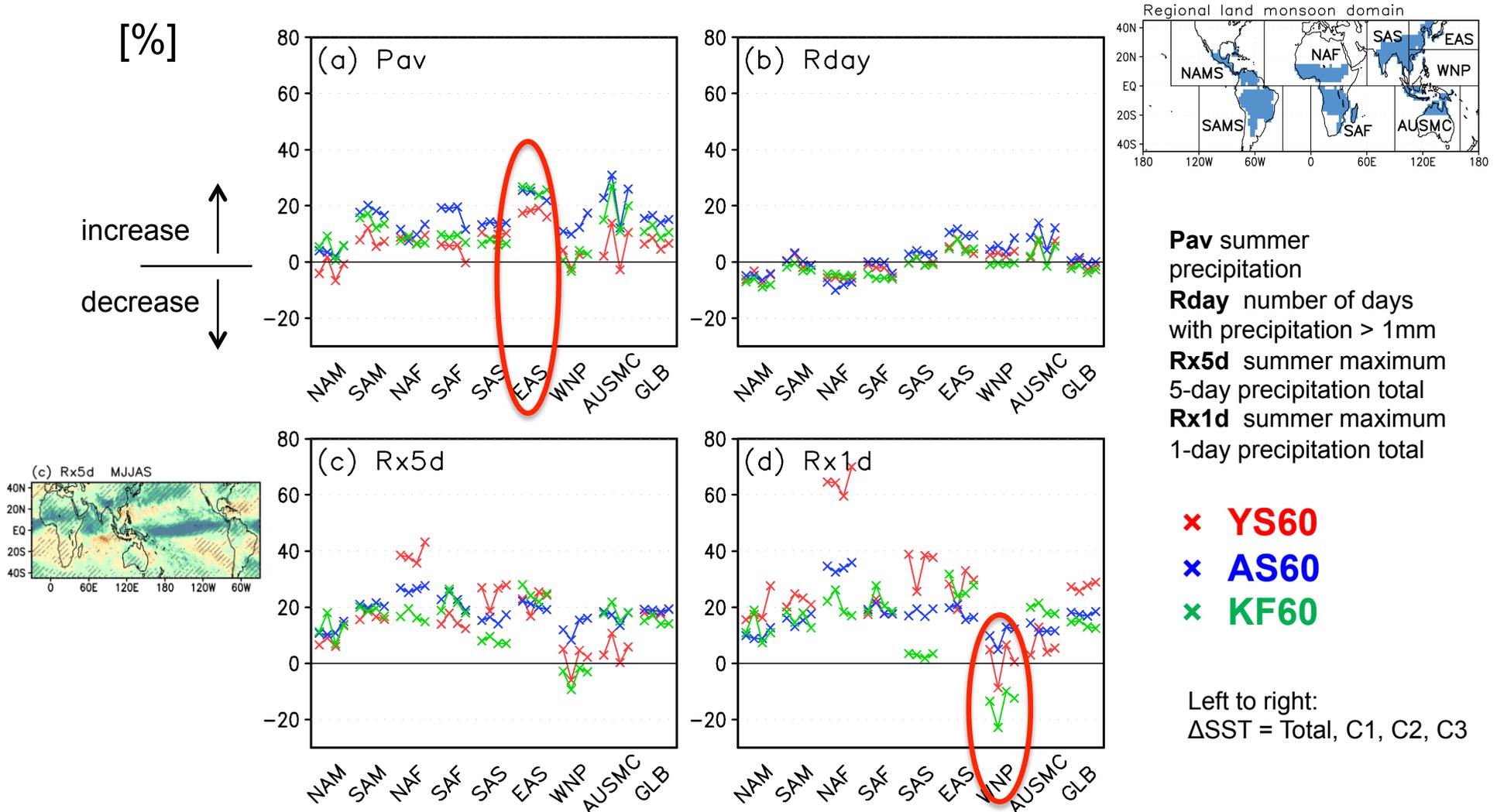
● 60kmAS

● 60kmKF

○ CMIP5 AGCMs

+ CMIP5 AOGCMs

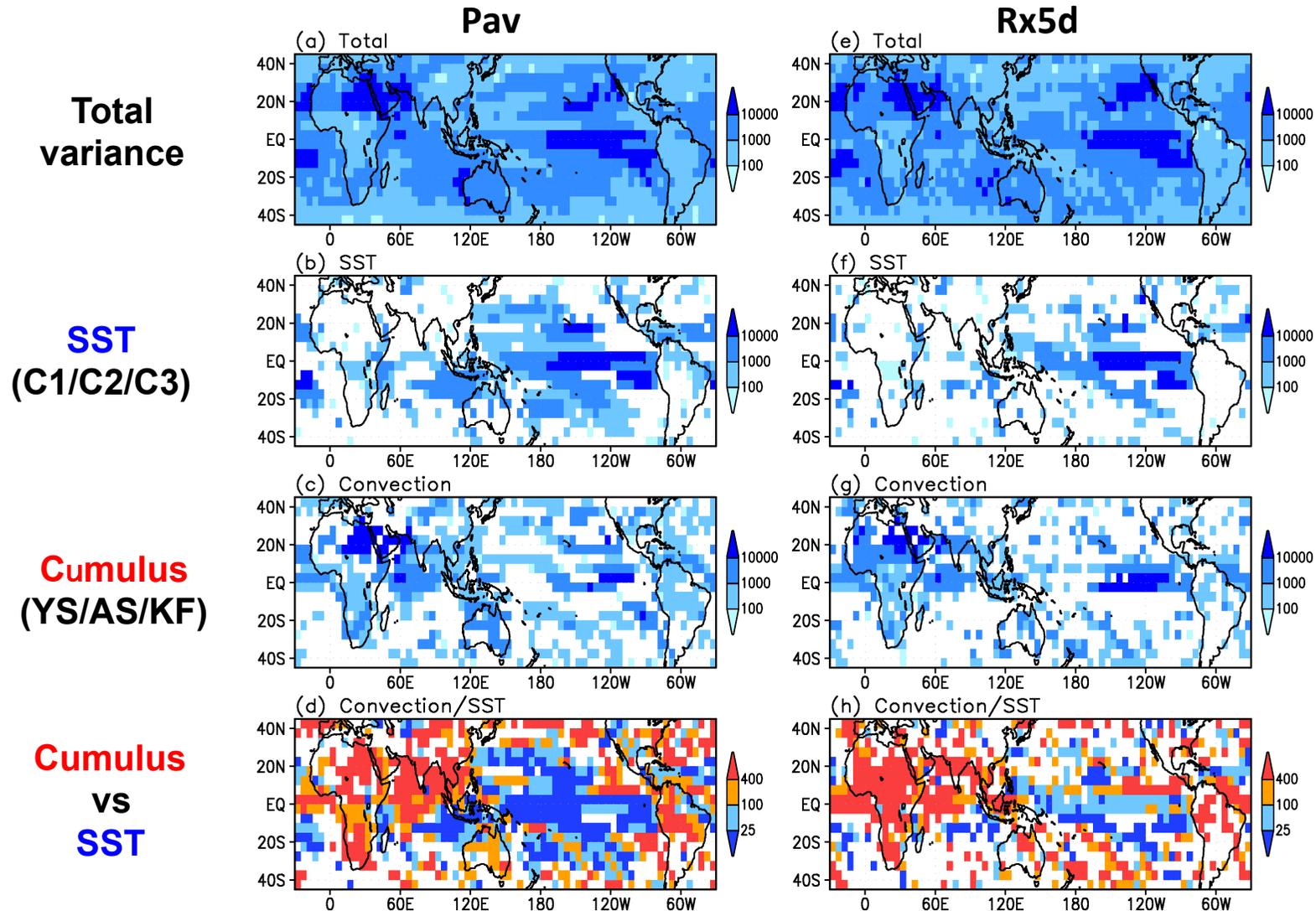
Future percentage changes in precipitation indices over the land monsoon regions



- **Pav**: increase in almost all regions.
Large increase rate in EAS. Large scatter in AUSMC. No change in NAM.
- **Rx5d/Rx1d**: increase in all regions except WNP. Larger increase rate than Pav.

Two-way ANOVA for summer season

MJJAS in NH
NDJFM in SH

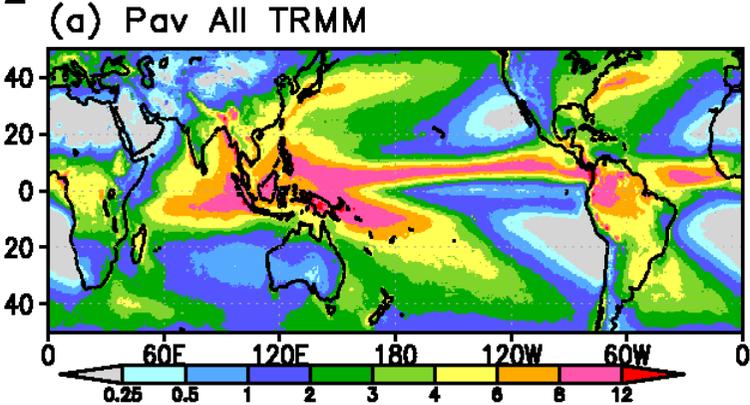


- **SST** is responsible for the uncertainty of precipitation amount over the tropical Pacific
> **SST change pattern in future is critical**
- **Cumulus** is responsible for the uncertainty of precipitation amount and extremes over most monsoon regions. More for extremes.

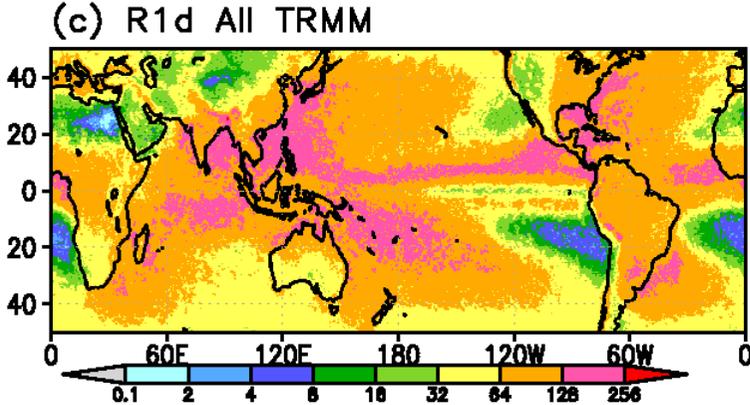
We investigate future changes in Rx1d and role of tropical cyclones (TCs)

TRMM3B42

Pav (total rainfall)

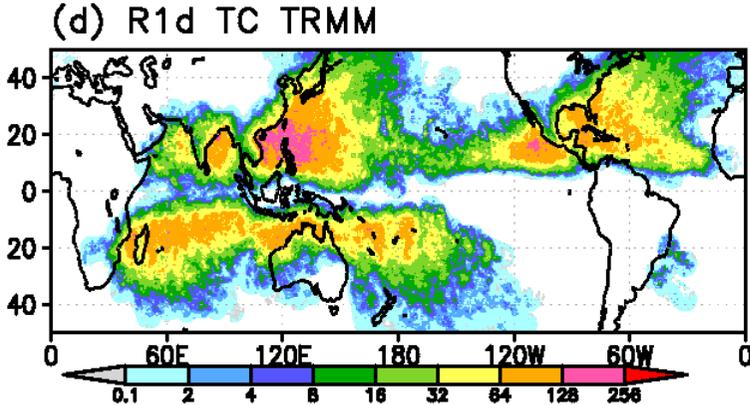
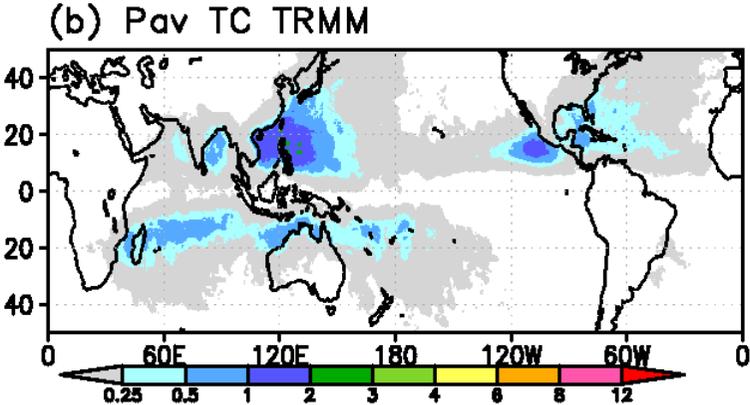


Rx1d (annual max)



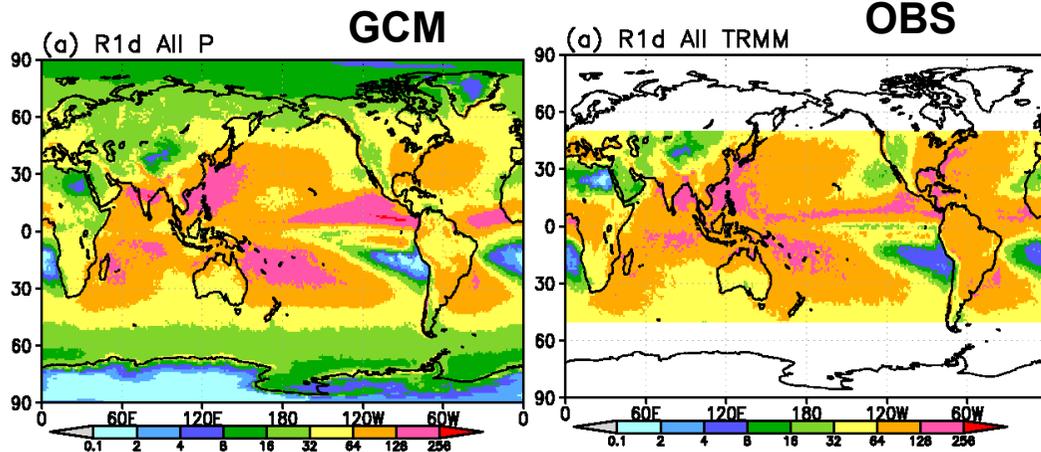
Extract rainfall within 500km from TC center

TC-rain



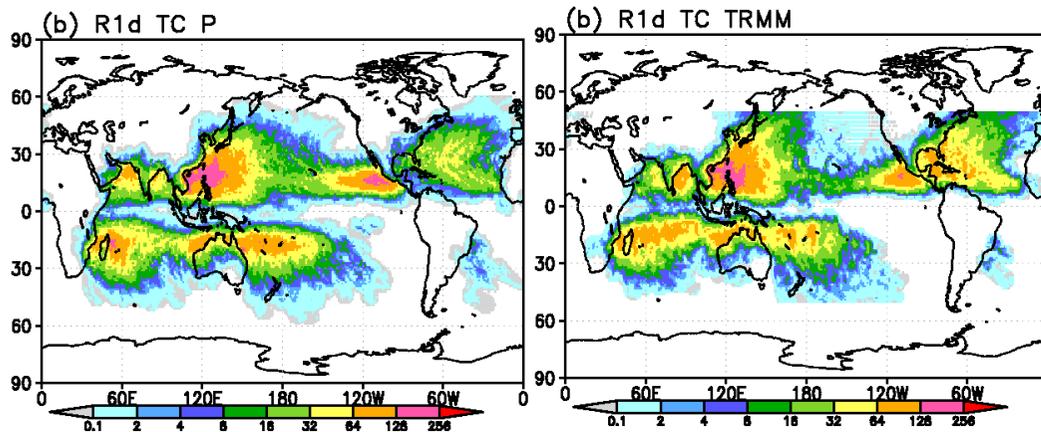
Rx1d associated with tropical cyclones

Rx1d



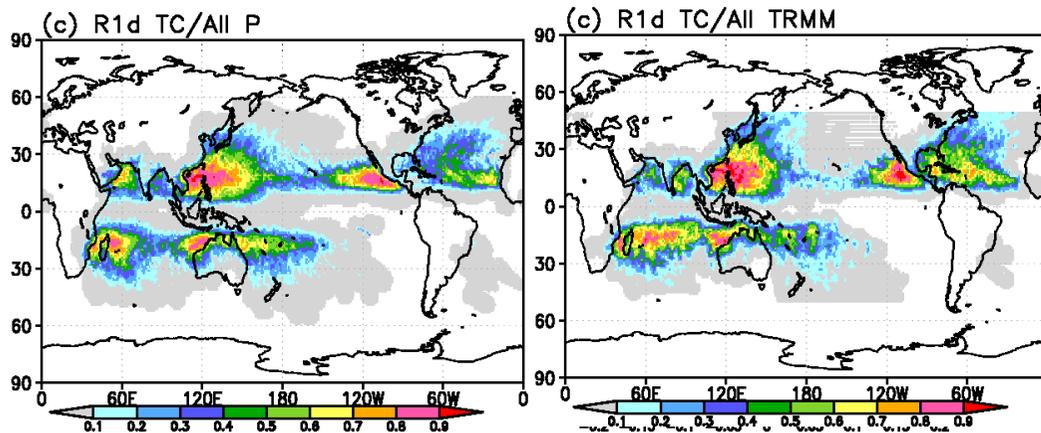
**Model
validation**

Rx1dTC



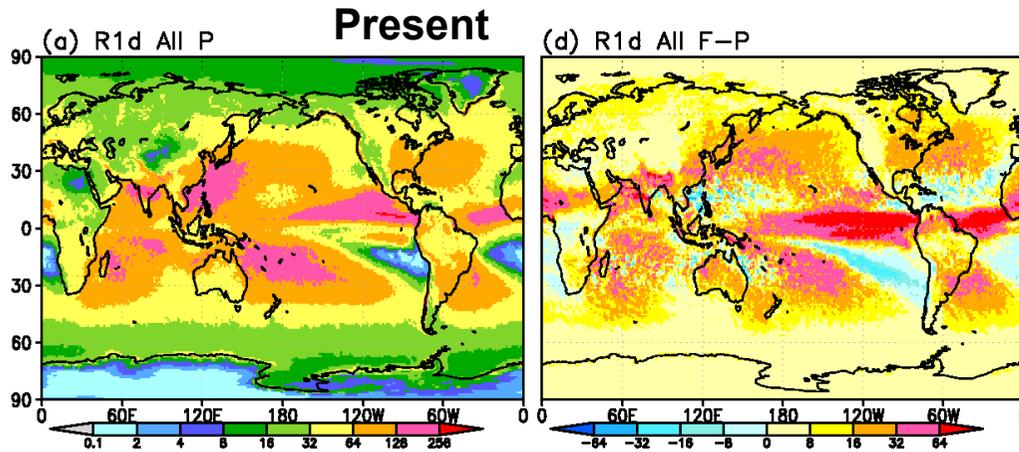
**20-km mesh AGCM
reproduces well
TC-associated
Rx1d,
quantitatively.**

Ratio
Rx1d-TC/Rx1d



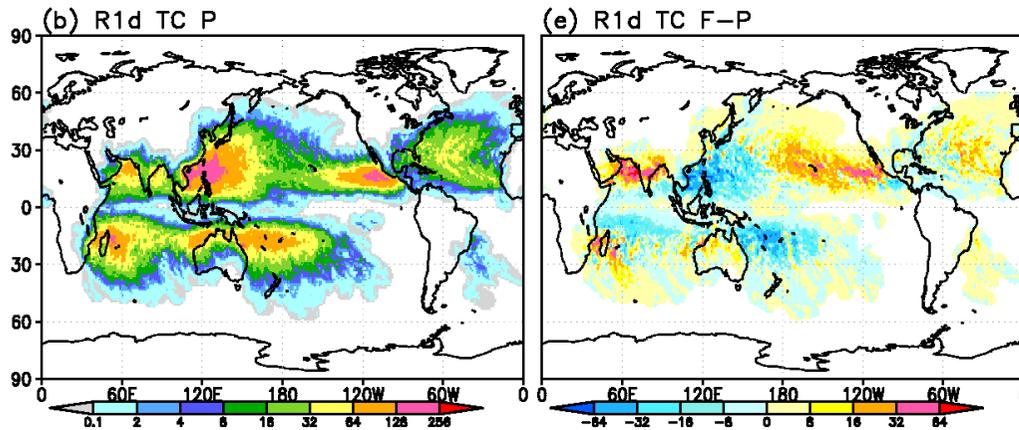
Rx1d associated with tropical cyclones

Rx1d



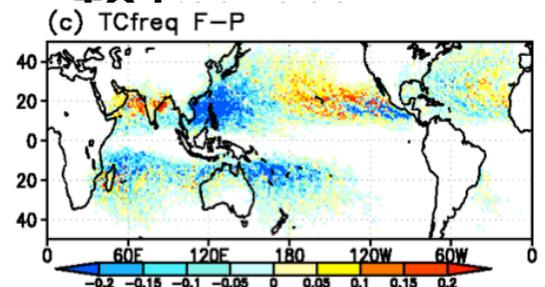
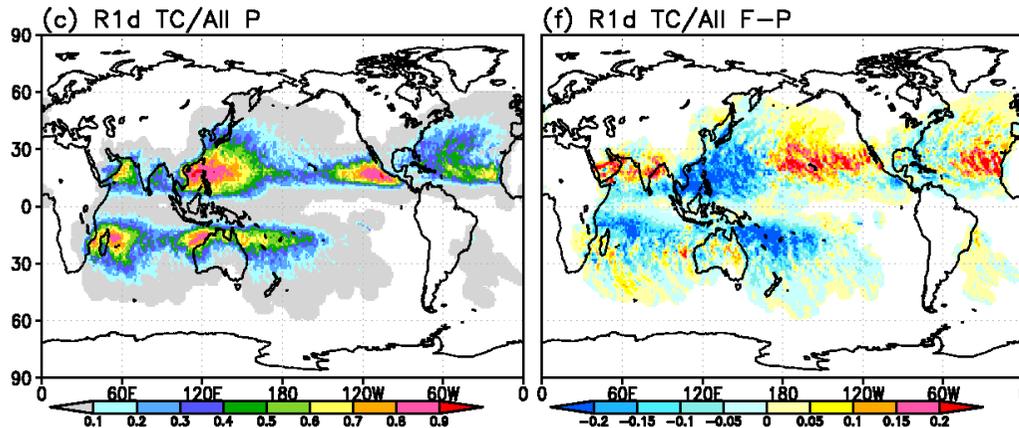
**Future
change**

Rx1d-TC



**TC-associated
Rx1d decreases in
the western
Pacific,
due to decreasing**

Ratio
Rx1d-TC/Rx1d



Kitoh and Endo (2016 SOLA)

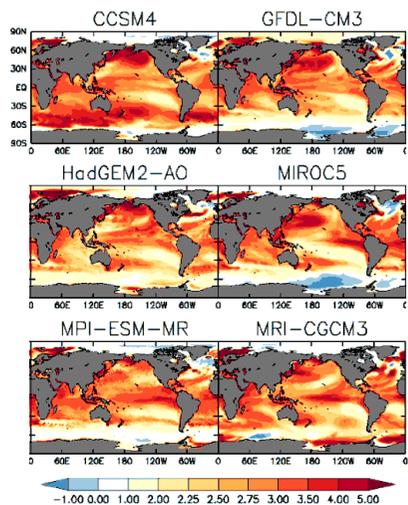
Large ensemble future climate simulation Database for Policy Decision-making for Future climate change (d4PDF)

6000 years for present and 5400 years for future
60-km mesh MRI-AGCM3.2

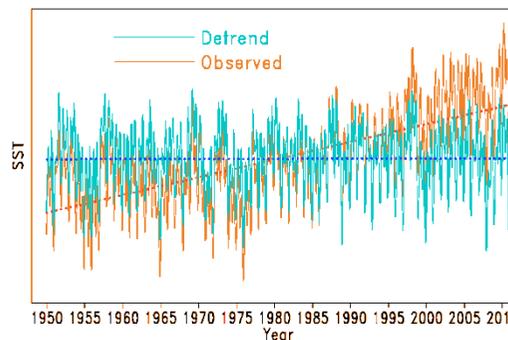


(Mizuta et al. 2017 BAMS)

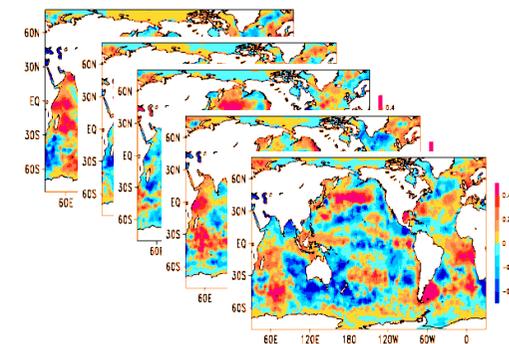
- present experiment
 - 60 years (1951-2010) x **100 members**
 - **4K warmer experiment**
 - 60 years under a 4K warmer climate than pre-industrial climate
 - SST : observed SST (detrend) + **CMIP5 AOGCMs ΔSST**
 - GHG/Aerosol/Ozone : 2090 year level based on the RCP8.5 scenario
 - **90 member ensemble**
 - 6 types of ΔSST patterns
 - 15 different atmospheric initial condition / SST perturbation
- 6×15 = 90 members



6 types of ΔSST pattern

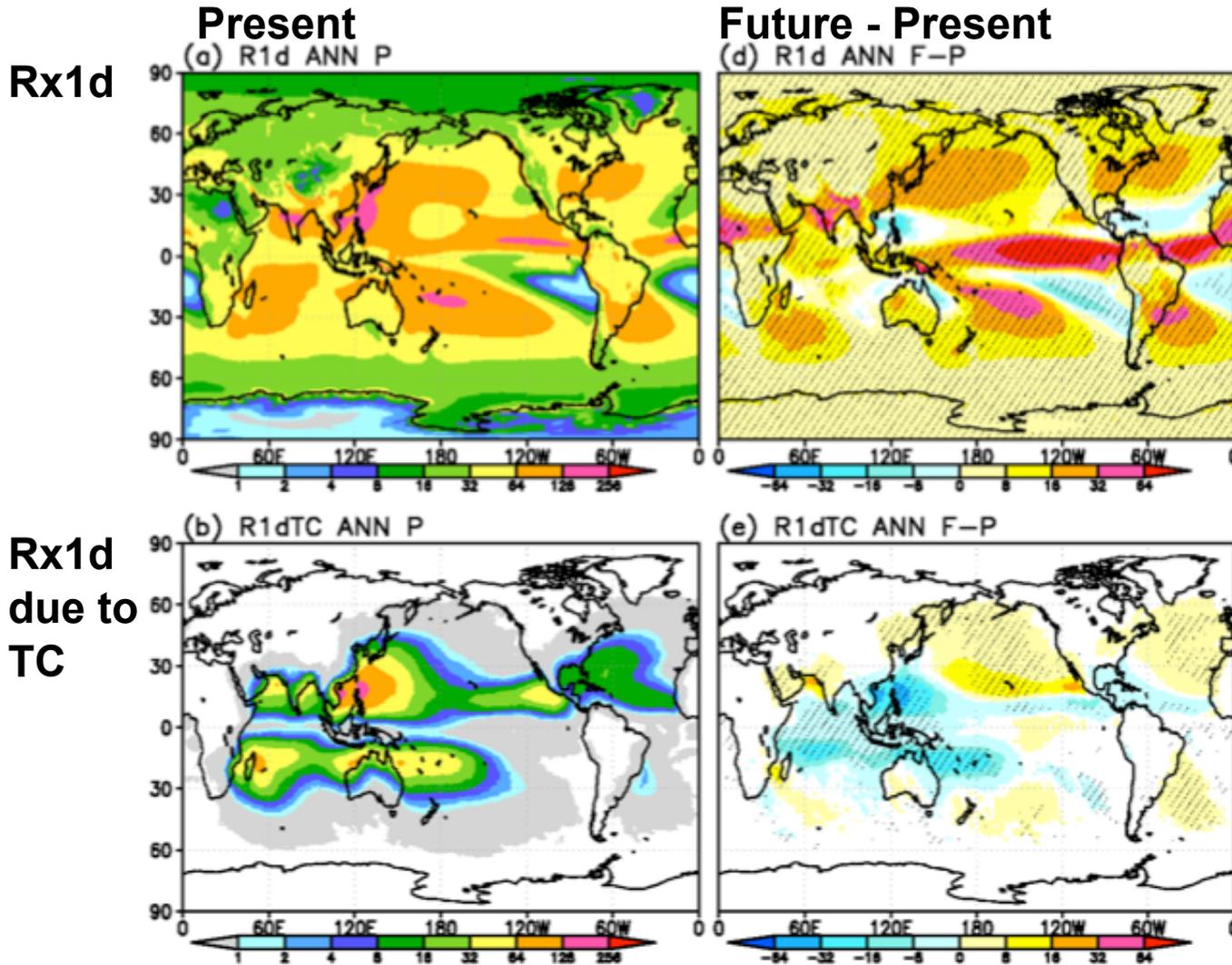


Observed SST (detrend)



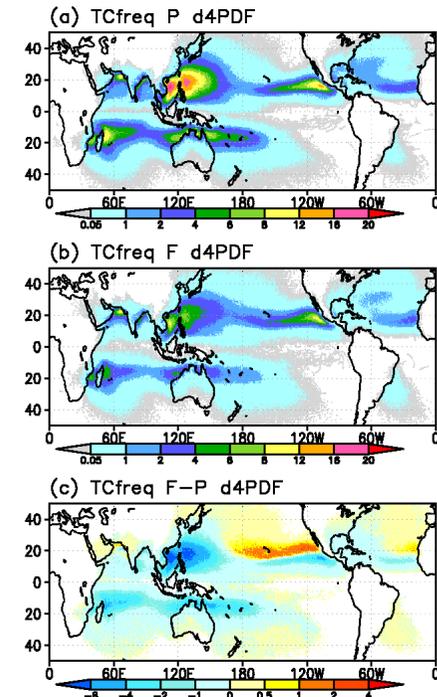
Different atmospheric initial condition / SST perturbations

Rx1d and tropical cyclone (TC)-associated Rx1d



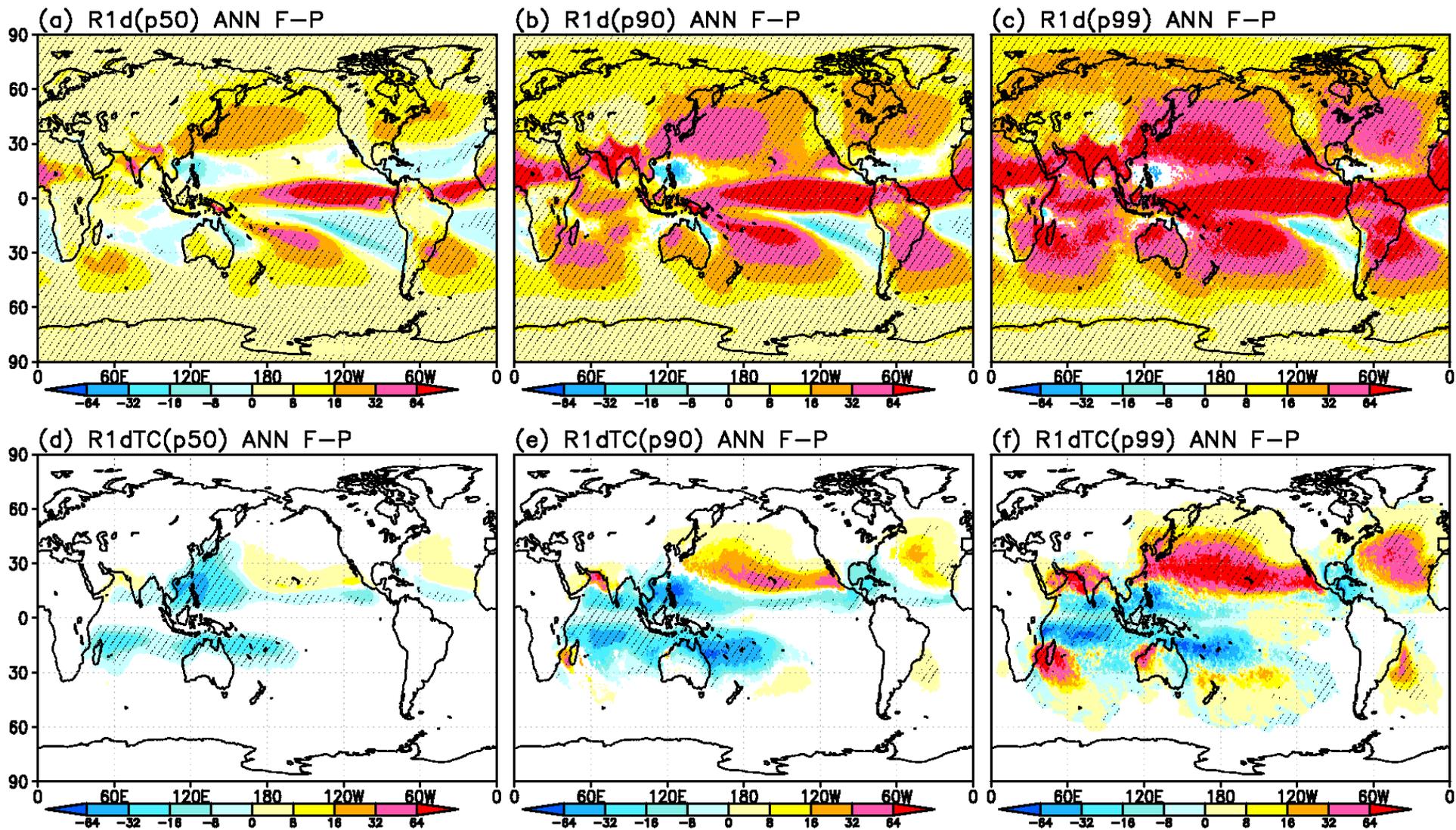
Hatch: same sign for all 6 Δ SST members

TC frequency



TC-associated Rx1d decrease in the western North Pacific due to TC frequency decrease

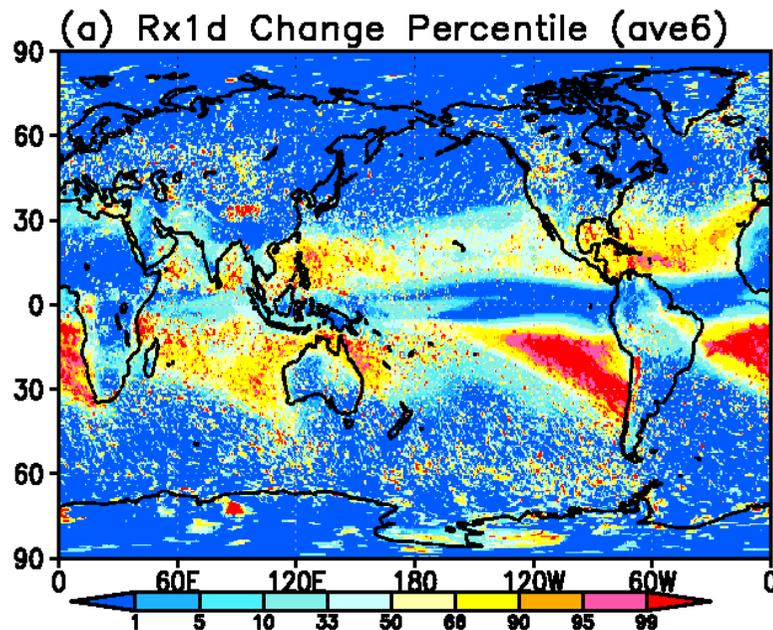
Future changes of (top) Rx1d and (bottom) TC-associated Rx1d (left) 50%-ile, (middle) 90%-ile, (right) 99%-ile



Color: 95% significant
Hatch: same sign for all 6 Δ SST members

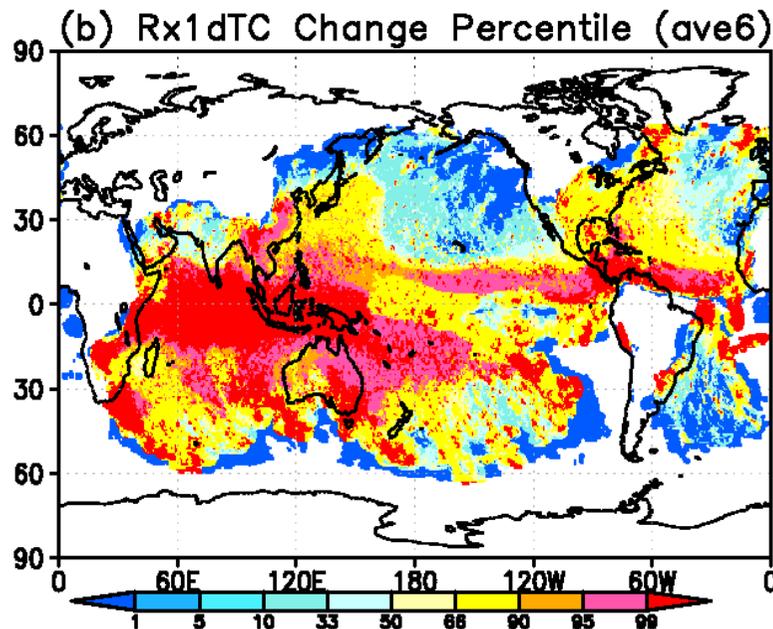
Over the northwestern tropical Pacific, median value of TC-associated Rx1d decreases, but its 90/99%-ile value increases around Japan

Percentile value at which future changes in (top) Rx1d and (bottom) TC-associated Rx1d become negative to positive



Over most of land, the threshold is below 33-percentile.

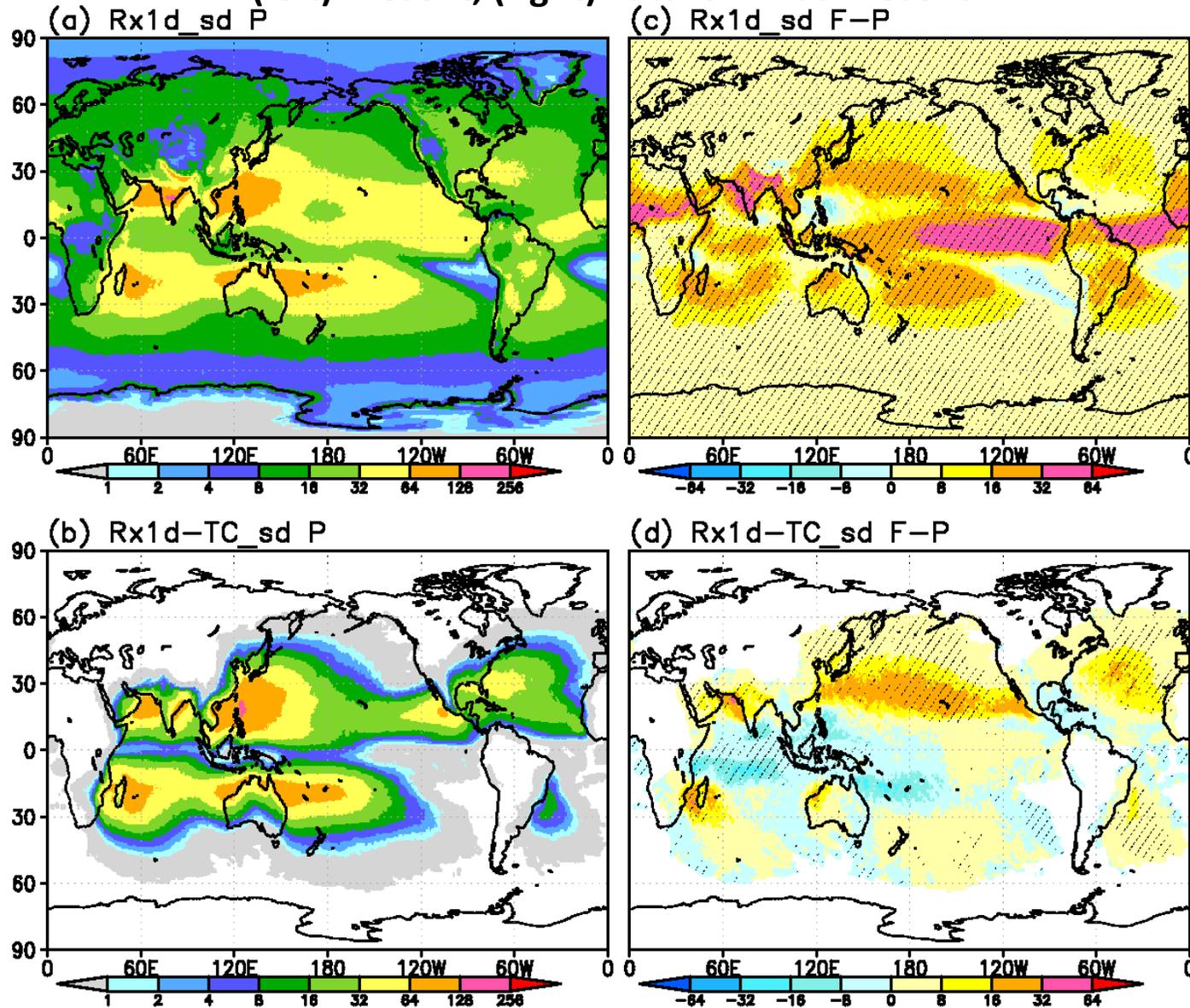
Rx1d above 90-percentile will increase over the area affected by TCs and over the subtropical oceanic dry areas.



Changes sign between 80- and 90-percentile in East Asia around Japan and eastern China.

Extreme TC-associated precipitation once in 5 years will become heavier in a future warmer climate.

Interannual standard deviation of (top) Rx1d and (bottom) TC-associated Rx1d
 (left) Present, (right) Future minus Present



Hatch: same sign for all 6 Δ SST members

Rx1d-TC increases in the NH subtropics from Hawaii to Japan

Summary

MRI-AGCM3.2 shows a high performance in simulating monsoon mean and extreme precipitation, with some dependency on the choice of convection scheme for extreme precipitation.

In a warmer climate, extreme precipitation (Rx1d) is projected to increase, except some regions around the western tropical Pacific. The latter is due to decreasing frequency of **tropical cyclones** (TC). However, interannual **variability and extremes** of Rx1d will increase.

Extreme Rx1d will increase more than the mean Rx1d. Sign of TC-associated Rx1d changes depends on the **percentile threshold**.

Interannual variability of Rx1d will increase in a region extending from Hawaii to the south Japan, implying an increasing risk of heavier rainfall events by global warming.

Role of TC is large in changes of precipitation extremes for some regions. Thus **model's ability in simulating TC is important**.