



the
abdus salam
international centre for theoretical physics

**COURSE ON CLIMATE VARIABILITY
STUDIES IN THE OCEAN
"Tracing & Modelling the Ocean Variability"
16 - 27 June 2003**

301/1507-10

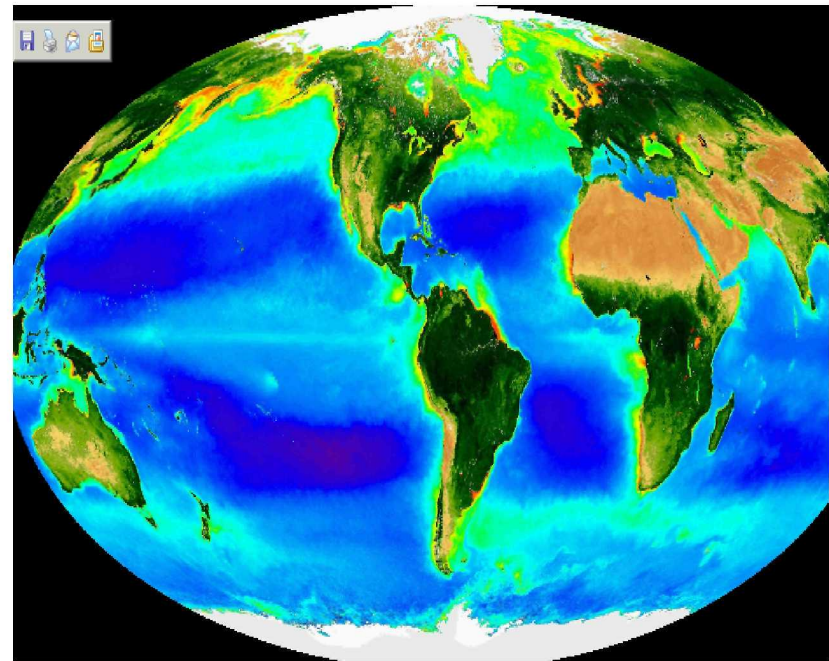
**Impact of physical processes on nutrients and
biological production**

**Ric G. Williams
University of Liverpool
United Kingdom**

Please note: These are preliminary notes intended for internal distribution only.

Impact of physical processes on nutrients and biological production

- How is biological production sustained?
- How are nutrient distributions controlled?



Williams & Follows (2003) Physical transport of nutrients and the maintenance of biological production. 19-51. In 'Ocean Biogeochemistry: The role of the ocean carbon cycle in global change'. Edited by Mike Fasham.

Seawifs average chlorophyll

Mar 21 to June 20 for 1998 to 2001
<http://seawifs.gsfc.nasa.gov/>

Impact of physical processes on nutrients and biological production

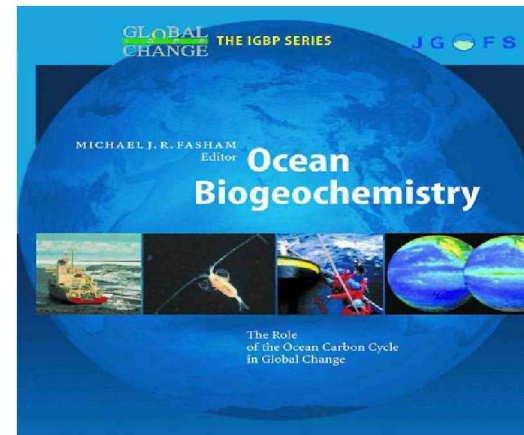
A. Mechanistic Review

1. Basin-scale distributions
2. Overturning
3. Convection
4. Gyre circulations
5. Eddies & Fronts
6. Interannual variability

B. How is biological production sustained in the Atlantic

Atlantic Meridional Transect Programme

Collaboration with Mick Follows (MIT)

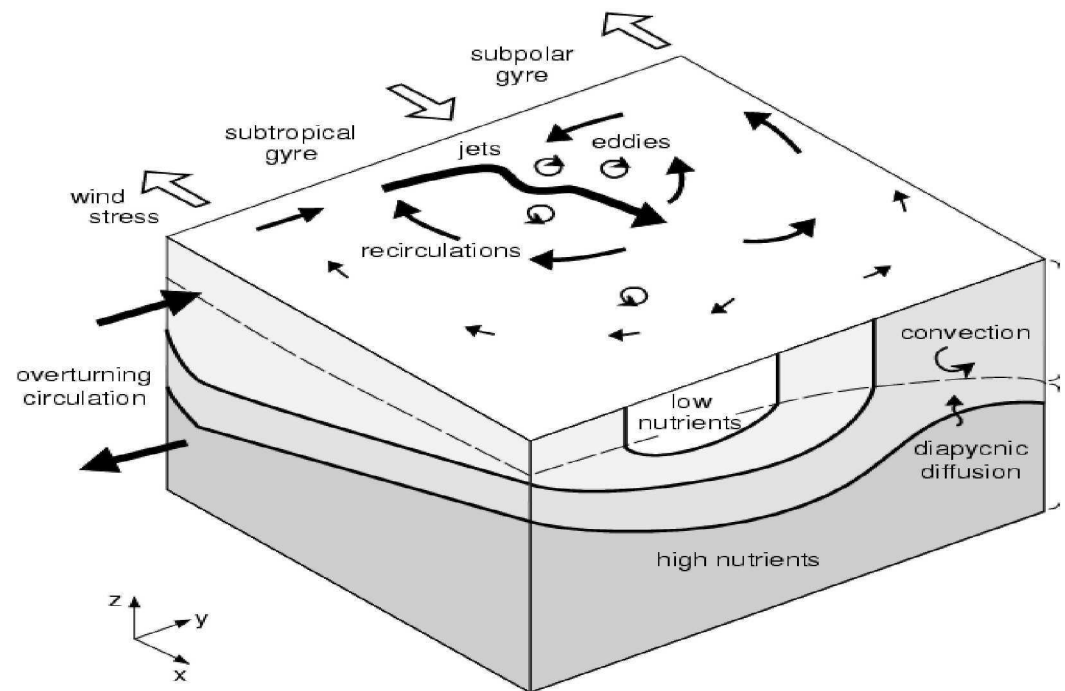
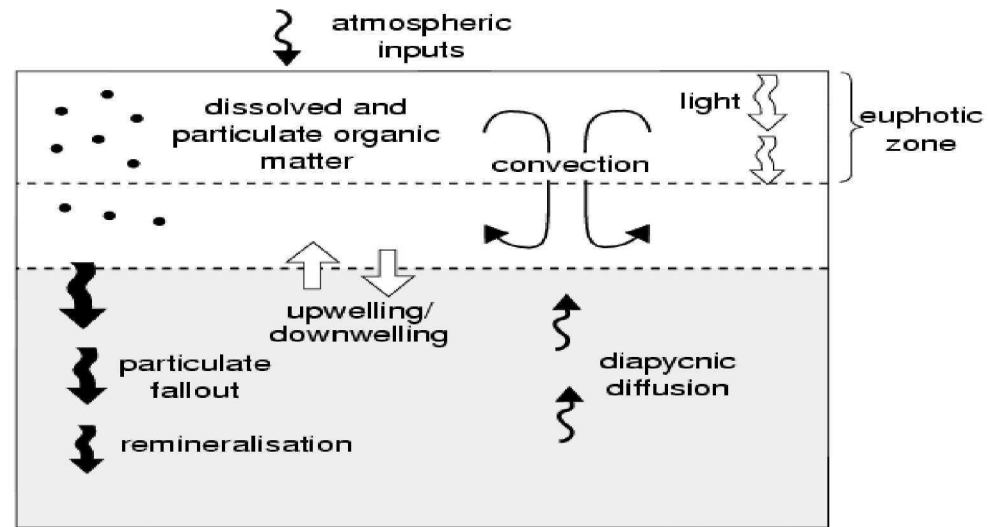


Williams & Follows (2003) In 'Ocean Biogeochemistry: The role of the ocean carbon cycle in global change'. Edited by Mike Fasham.

Collaboration with Claire Mahaffey & George Wolff (Liverpool) + AMT partners

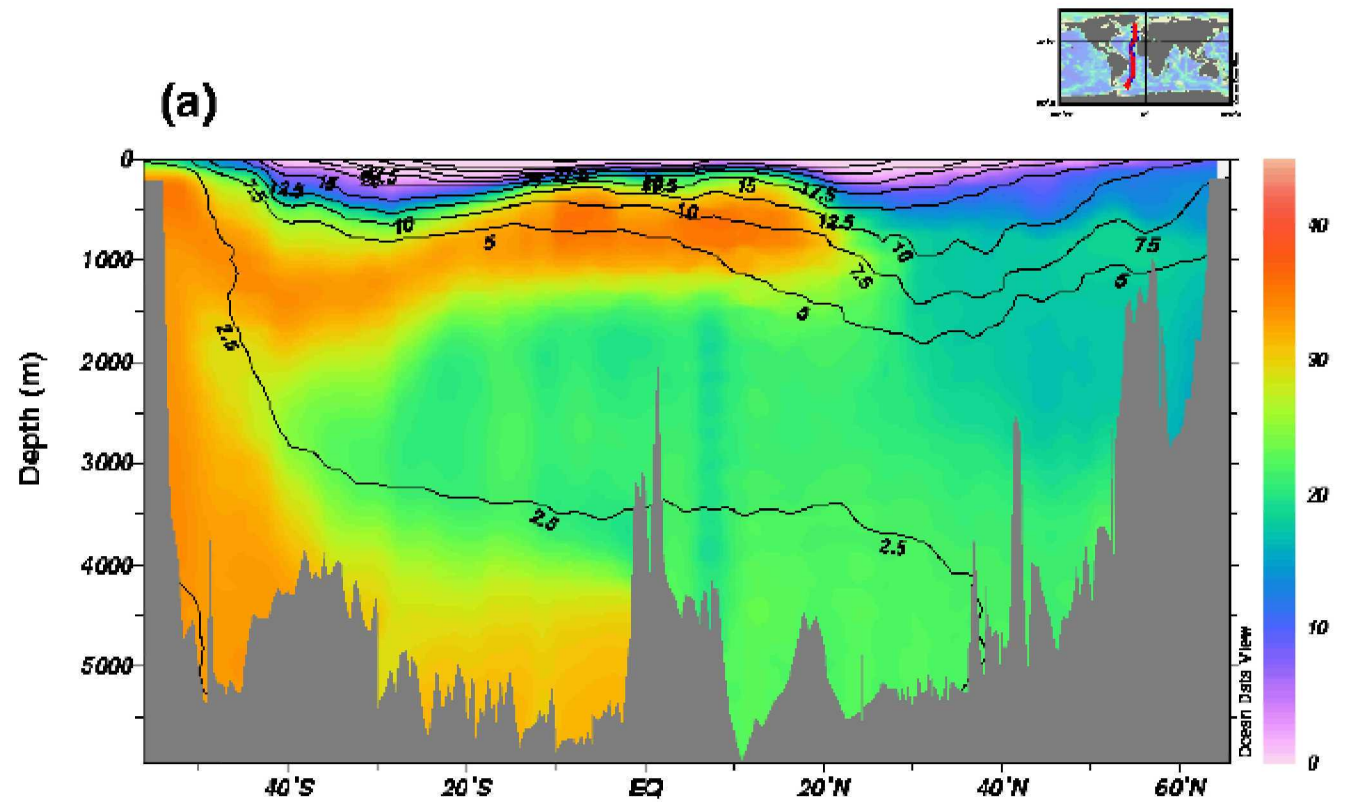
1. Basics

- Biological growth is in sunlit, surface ocean
- Organic matter formed with C : N : P ratio 112 : 16 : 1
- Gravitational fallout requires nutrients to be supplied to surface



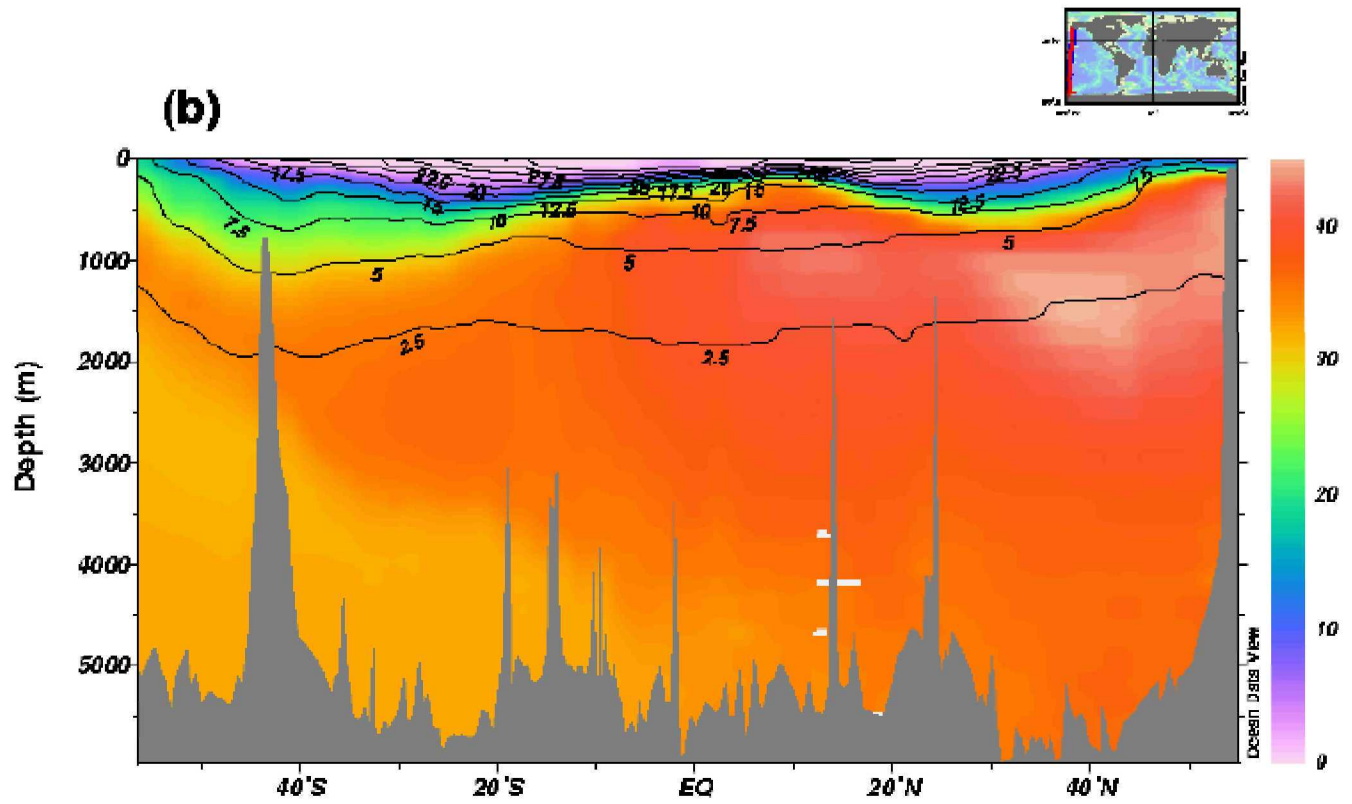
Atlantic

NO_3 ($\mu\text{mol kg}^{-1}$)



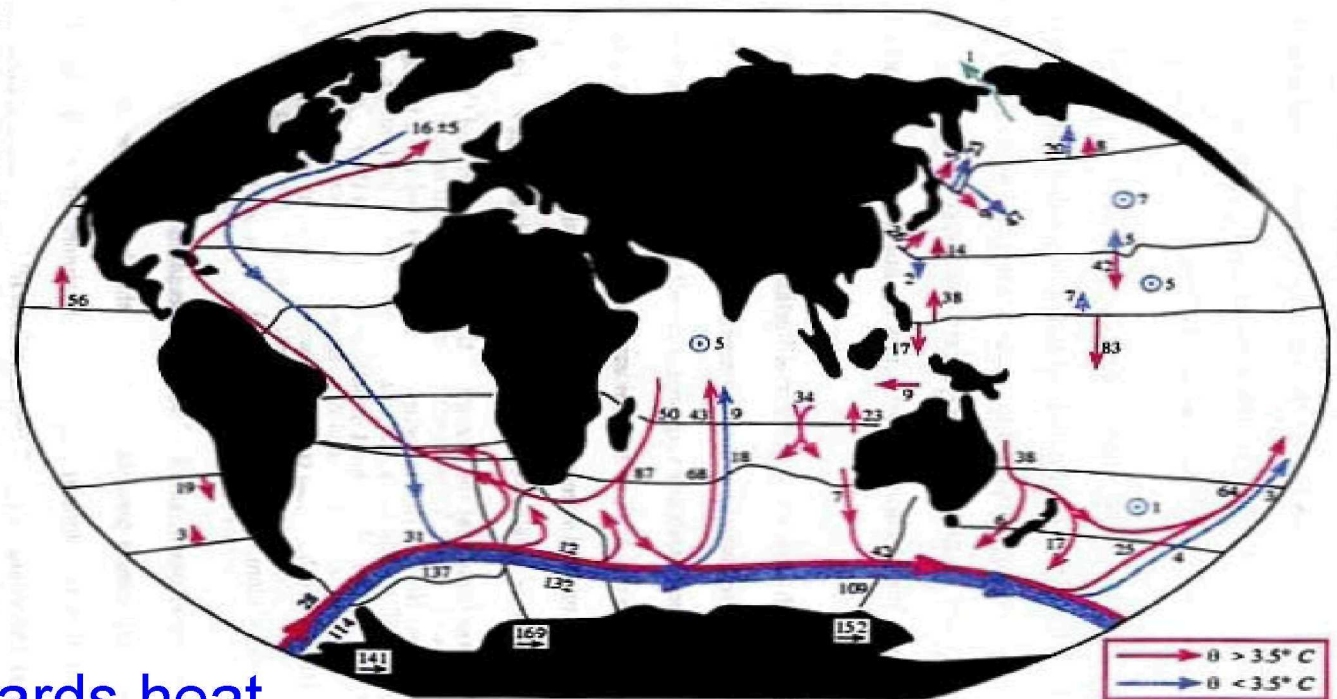
WOCE A16

Pacific



WOCE P15

2. Global circulation



- Atlantic northwards heat flux

- S. Ocean connects basins

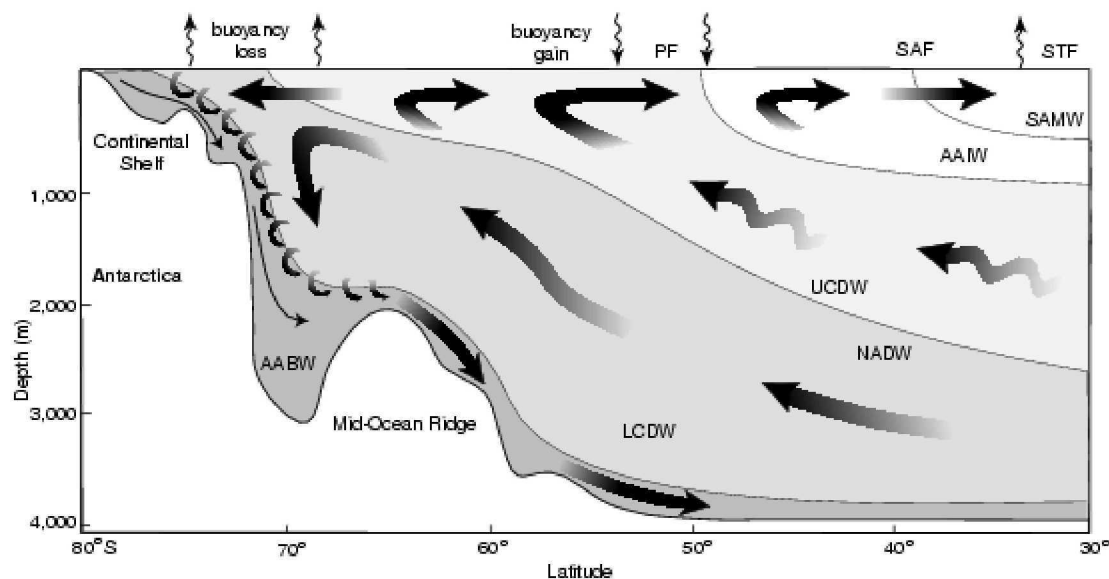
- Recirculations

Macdonald and Wunsch (1996) inversion

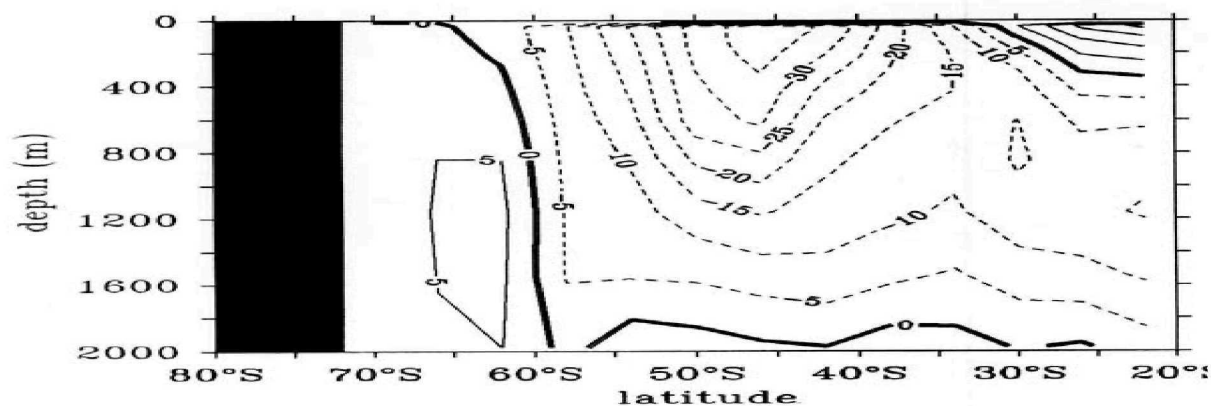
Volume flux of **warm** & **cold** water

Southern Ocean

- northwards Ekman flux
- northwards bottom flux
- eddy-driven southwards flux at mid-depths



Speer et al. (2000)



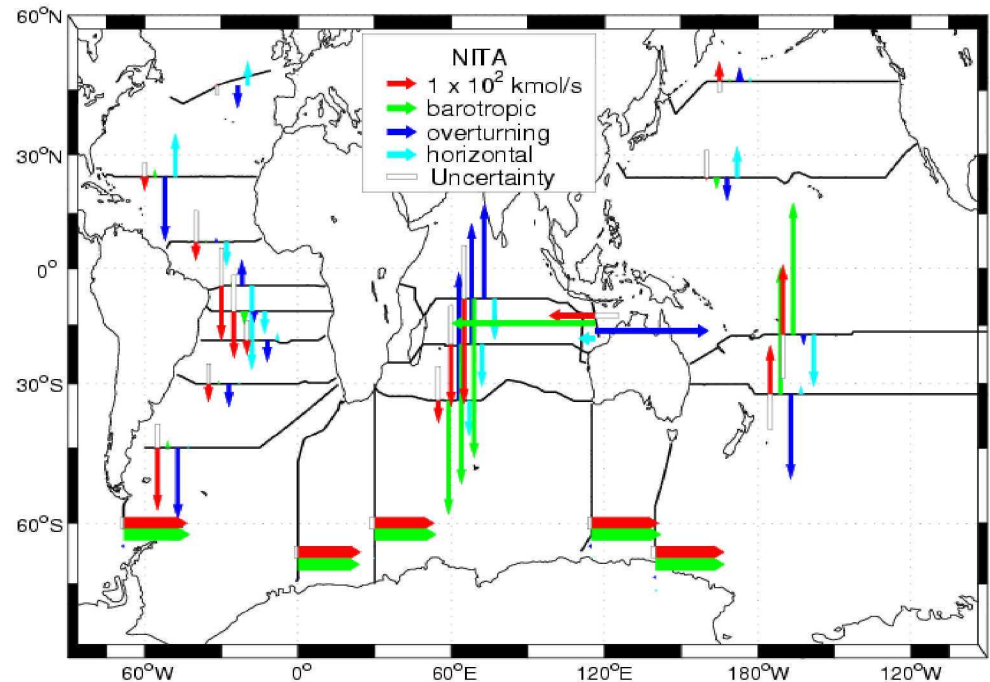
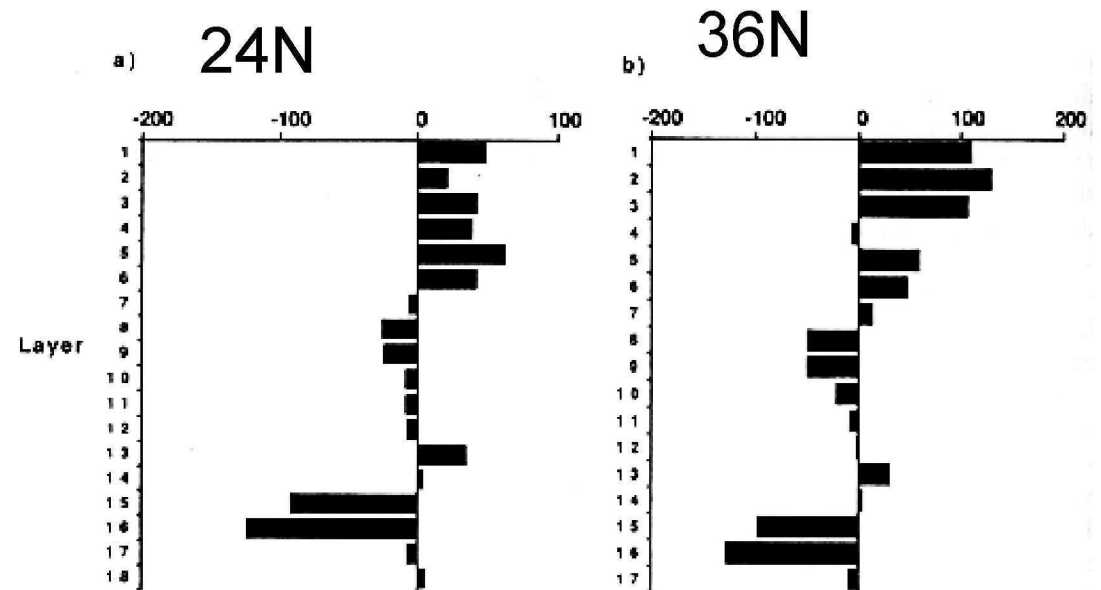
*Typical time-mean overturning streamfunction
if eddy effects not included*

Nitrate transport

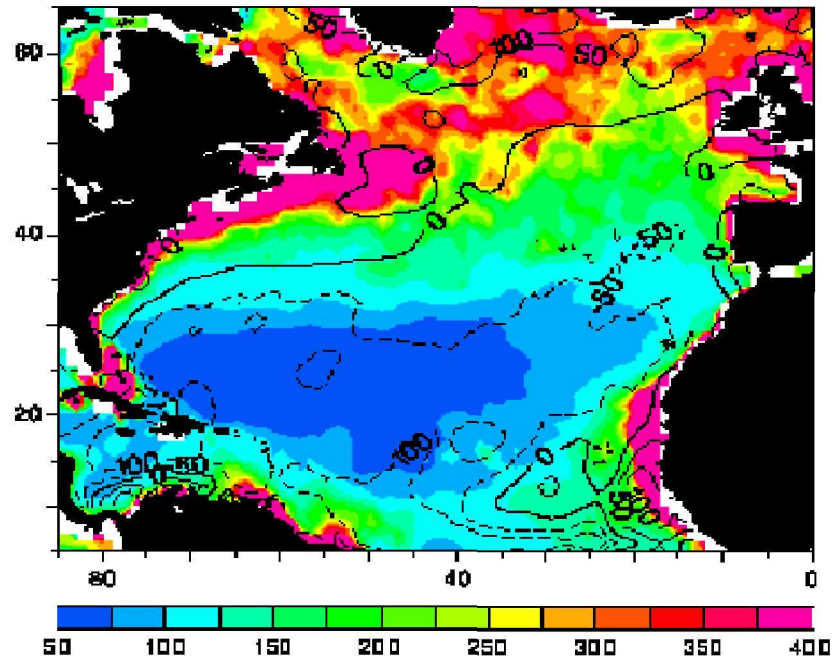
*Rintoul & Wunsch (1991)
for N. Atlantic*

- predicts northwards NO_3 flux at 36N
- loss of NO_3 between 24N and 36N

*Ganachaud & Wunsch
(2002) global inversion*

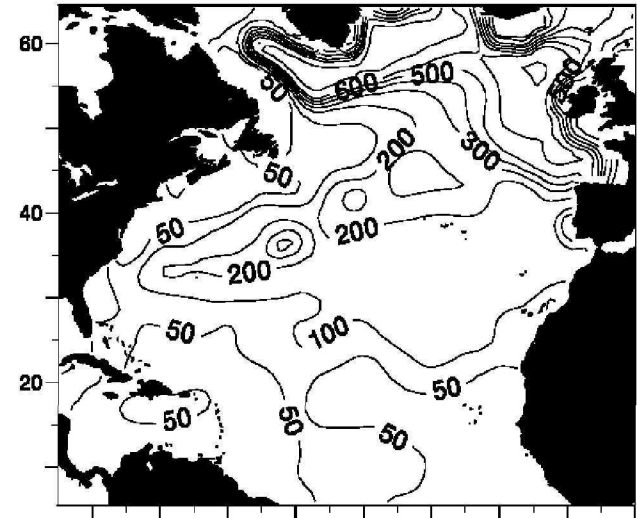


3. Convection

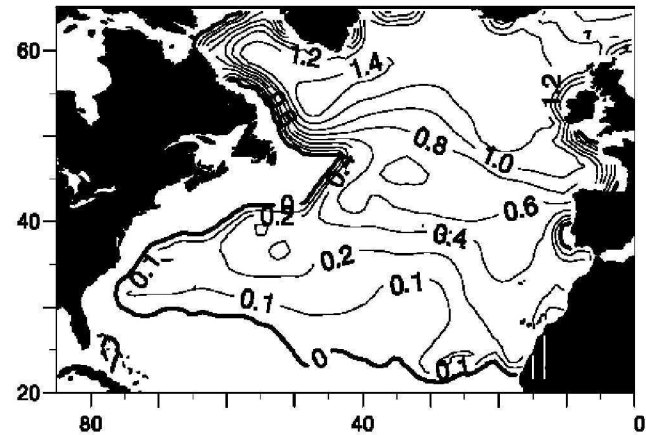


Satellite estimate of export production (gC m⁻² y⁻¹)
Sathyendranathan et al. (1995)

b) March mixed-layer thickness



c) Climatological convective nitrate flux

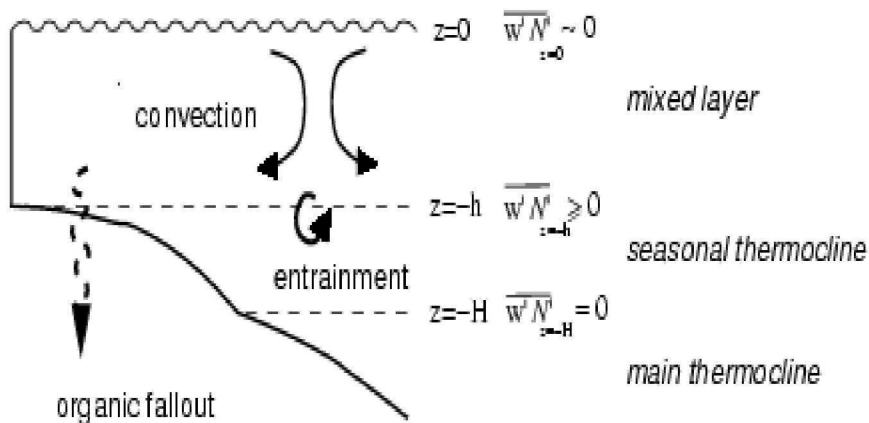
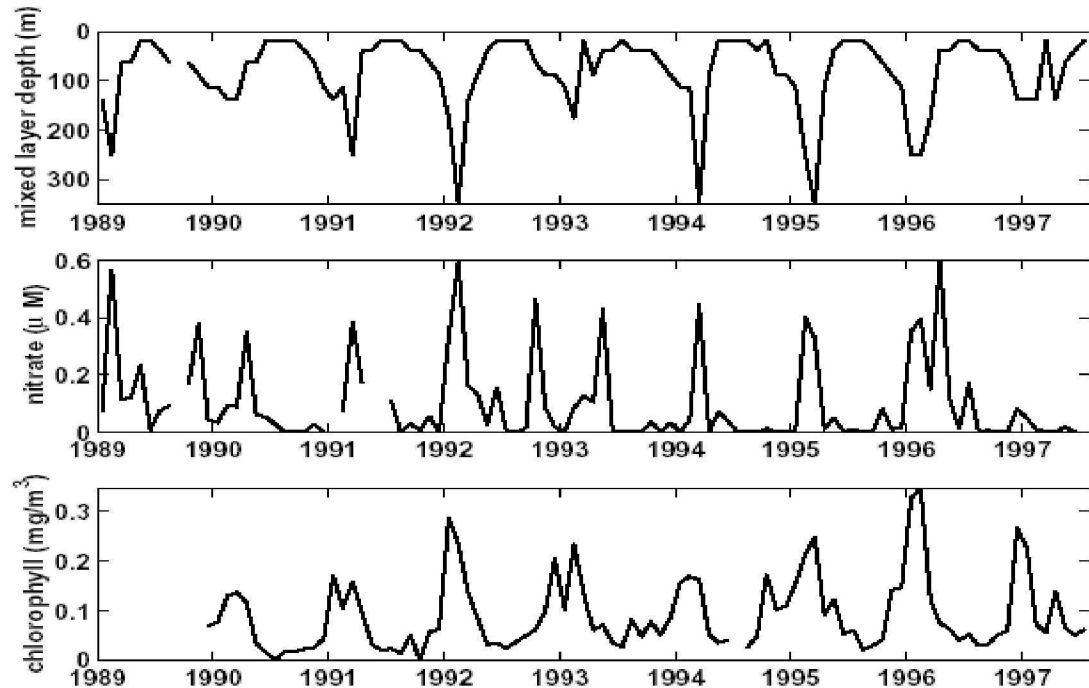


Williams & Follows (2003)

Is convection alone sufficient?

Convective variability

Variability at Bermuda



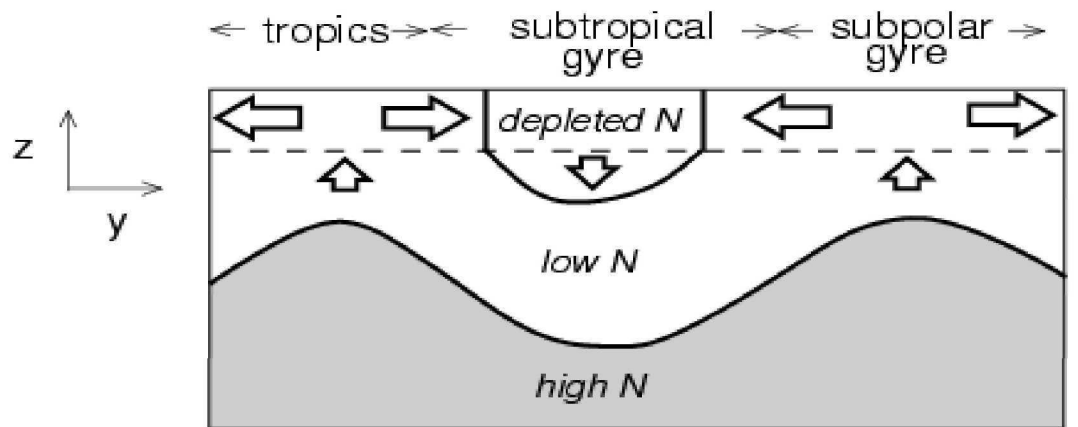
Eventually need a N source above main thermocline

— otherwise N diminish

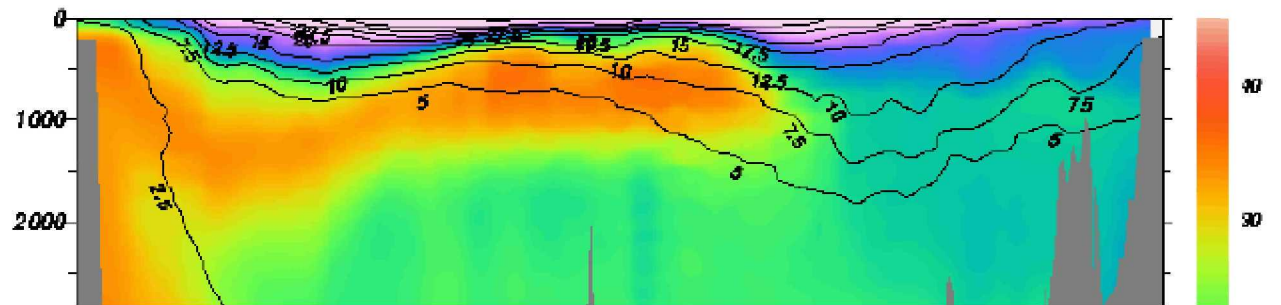
4. Gyres

*Characteristic
thermocline
undulations*

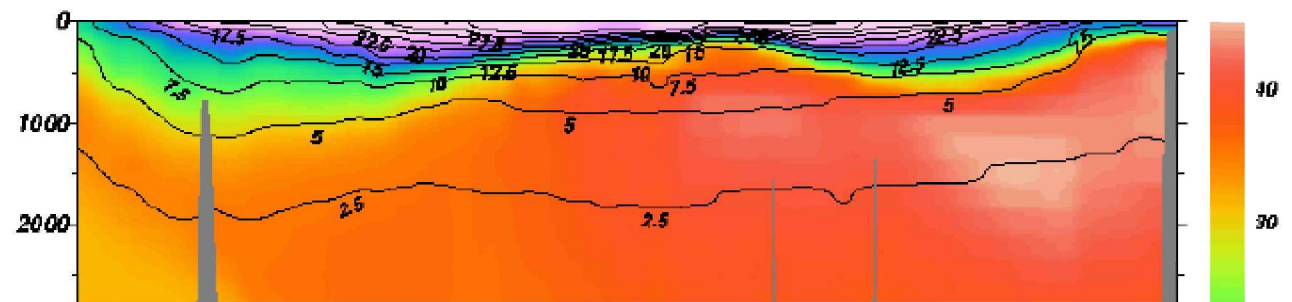
a) Ekman volume flux over a basin



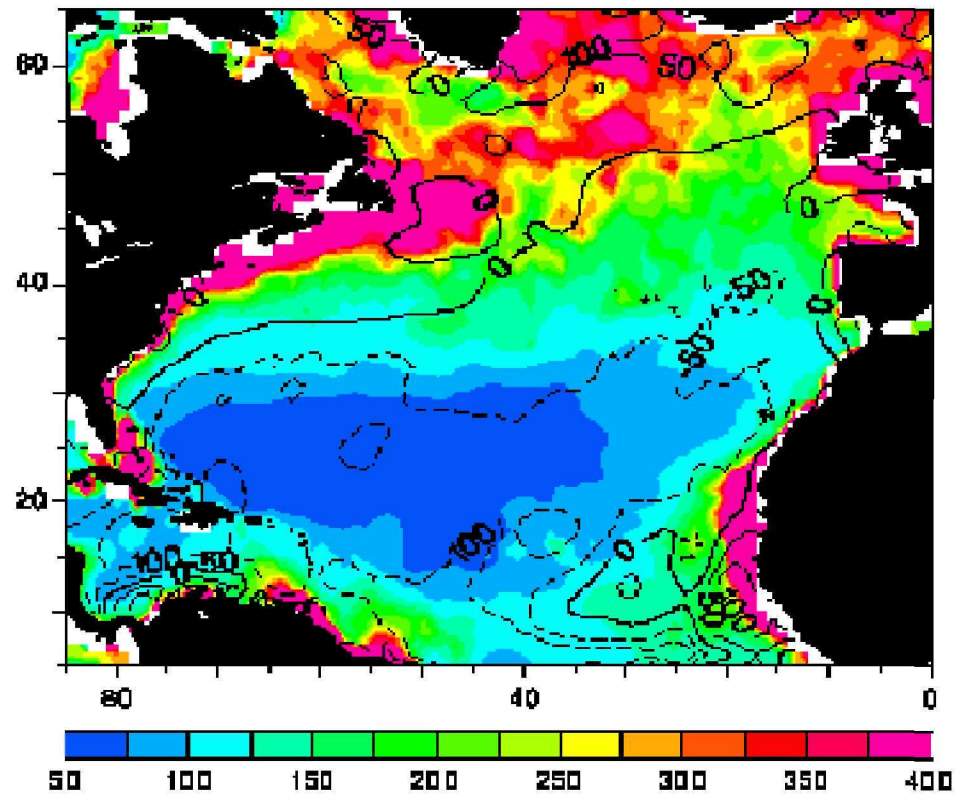
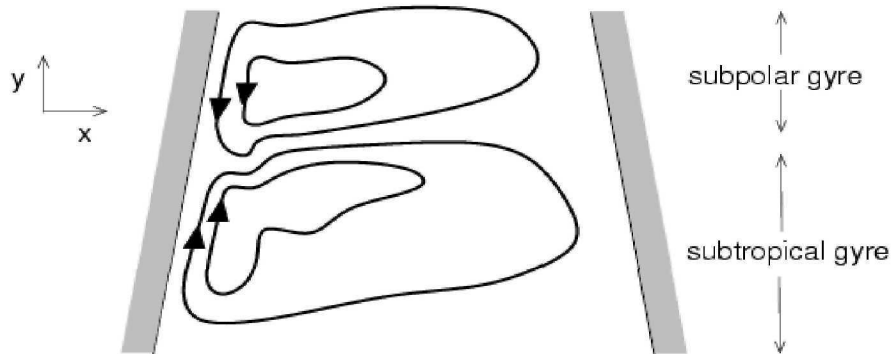
*Atlantic NO_3
section*



*Pacific NO_3
section*



Gyres

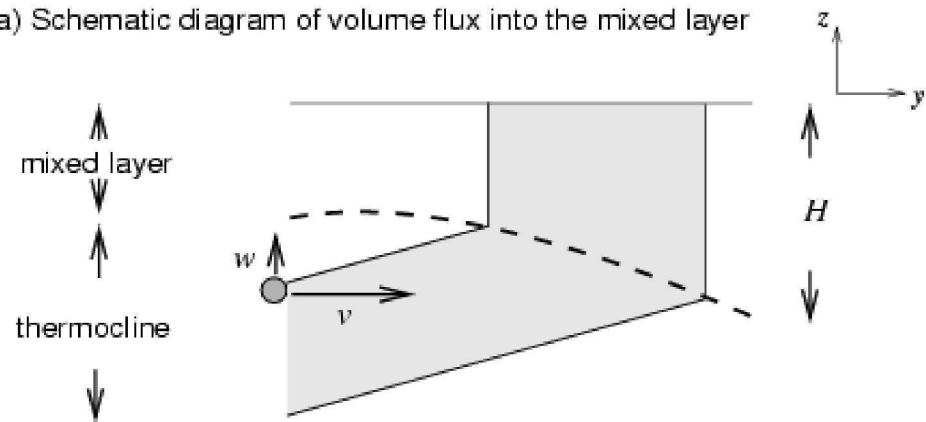


w (m y^{-1}) (contours)
Export production
($\text{gC m}^{-2} \text{y}^{-1}$) (colour)

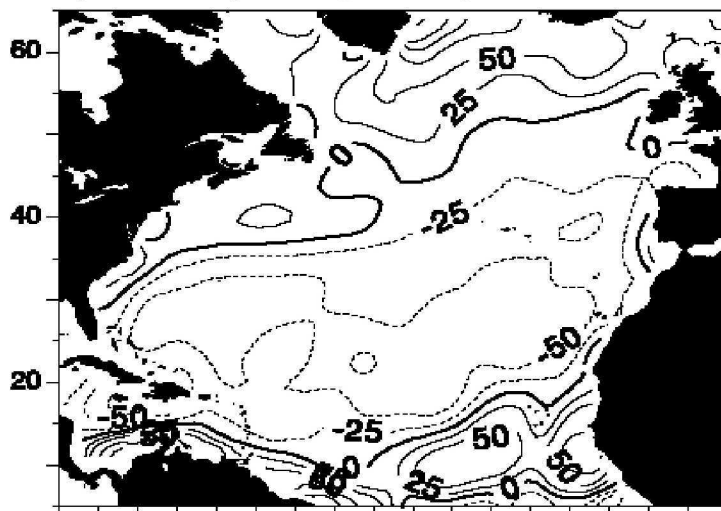
Does the vertical velocity control production?

Subduction

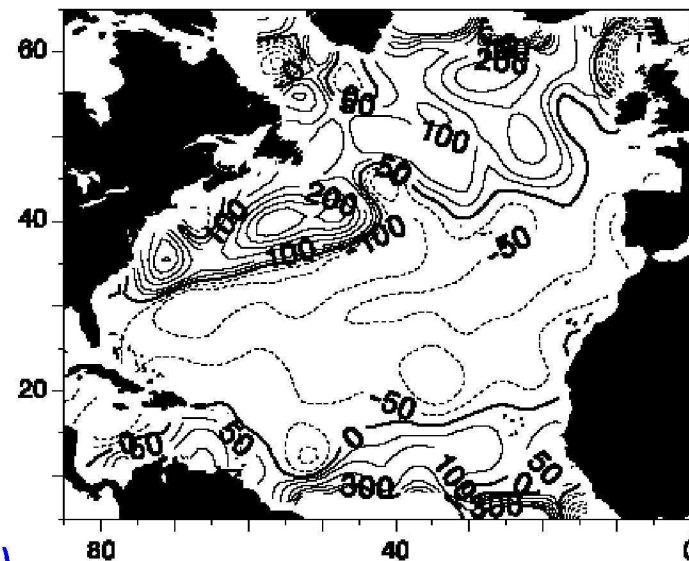
a) Schematic diagram of volume flux into the mixed layer



b) Ekman upwelling velocity



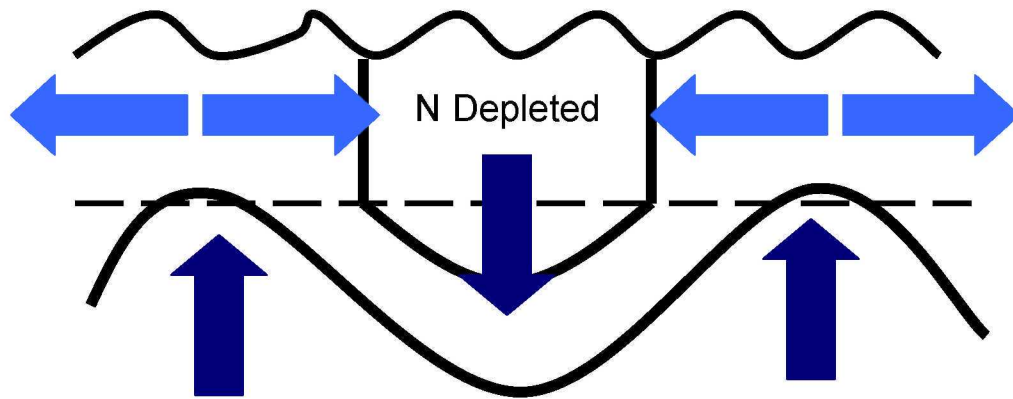
c) Volume flux into seasonal-boundary layer



Marshall, Nurser & Williams (1993)

*Lateral influx into mixed layer in subpolar gyre
— expect high influx of NO_3*

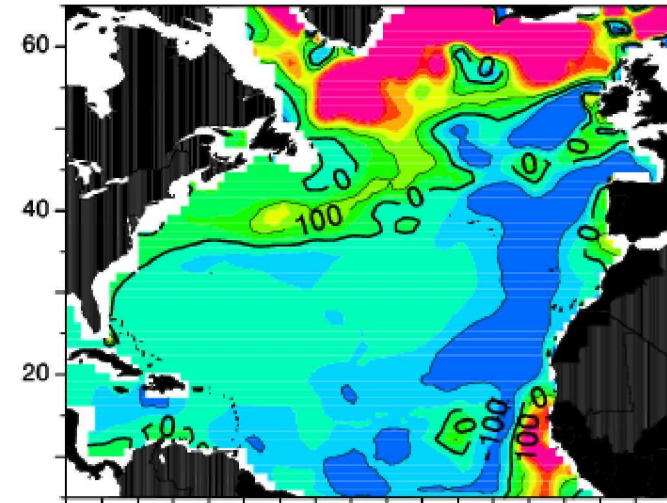
Lateral transfer of nitrate



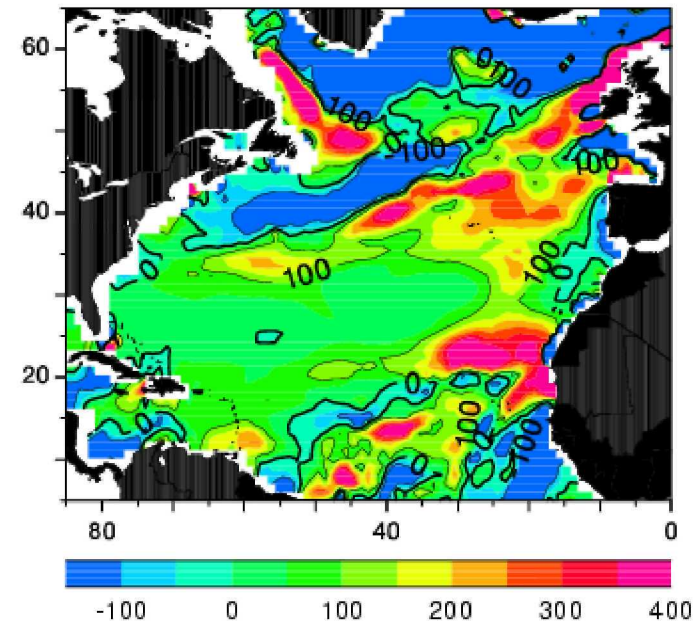
How does NO_3 get back to the surface?

Williams & Follows(1998)

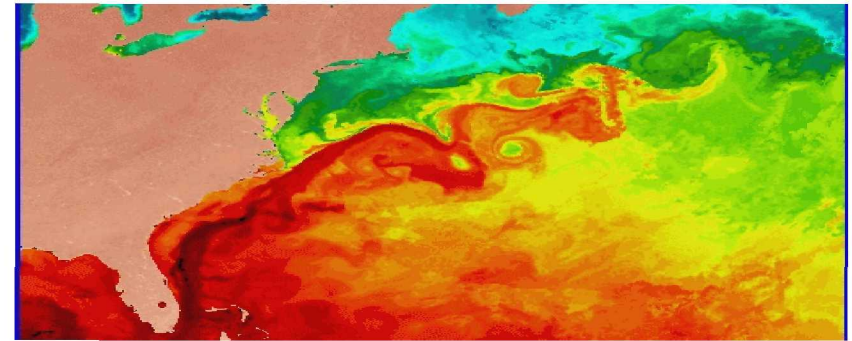
a) vertical Ekman nitrate flux



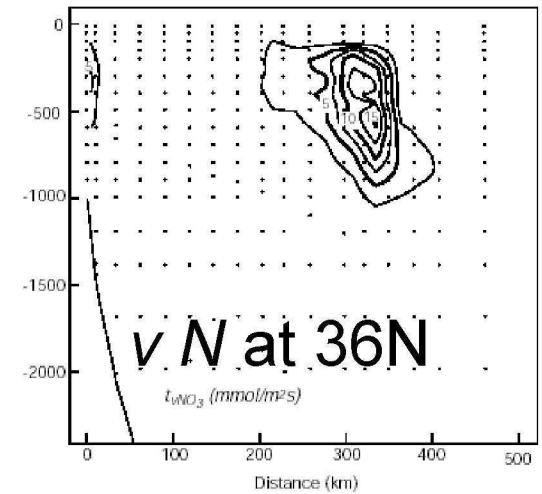
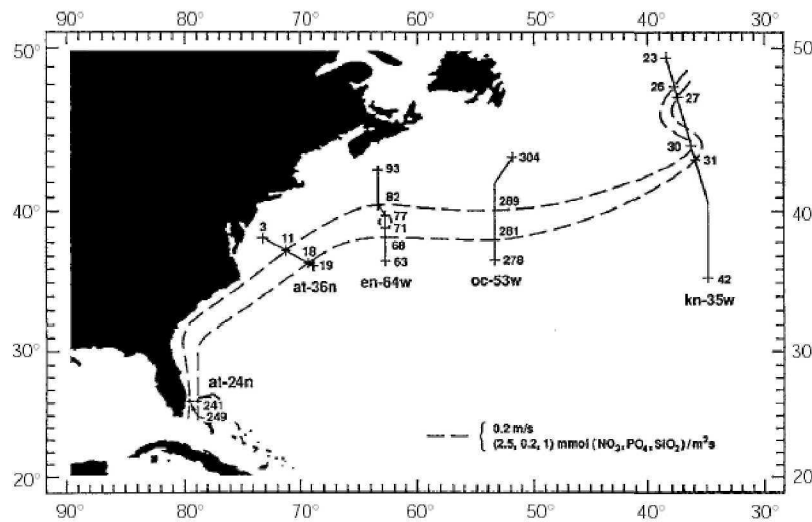
b) horizontal Ekman nitrate flux



Nutrient streams

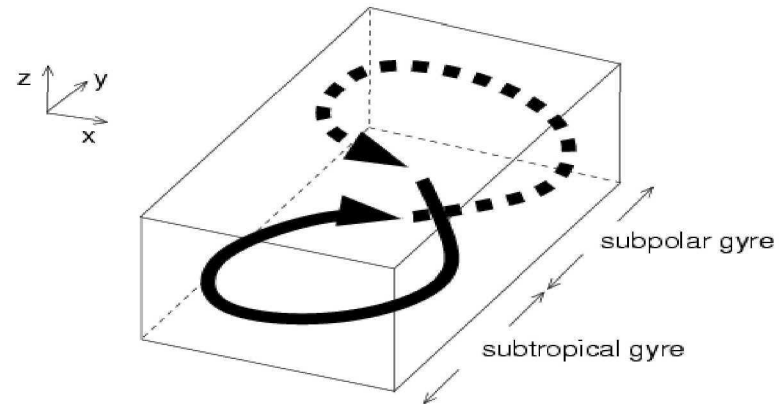


'Nutrient Stream' in the North Atlantic



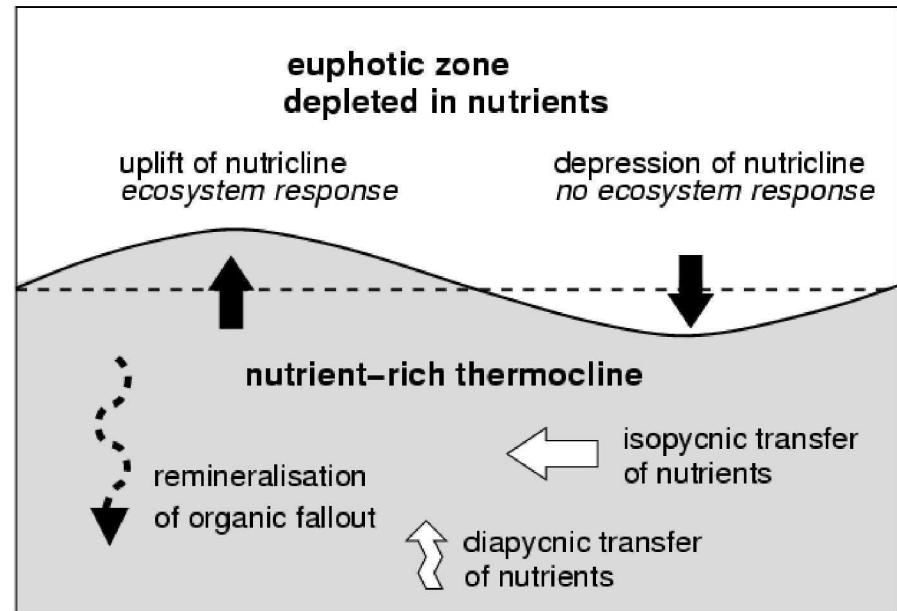
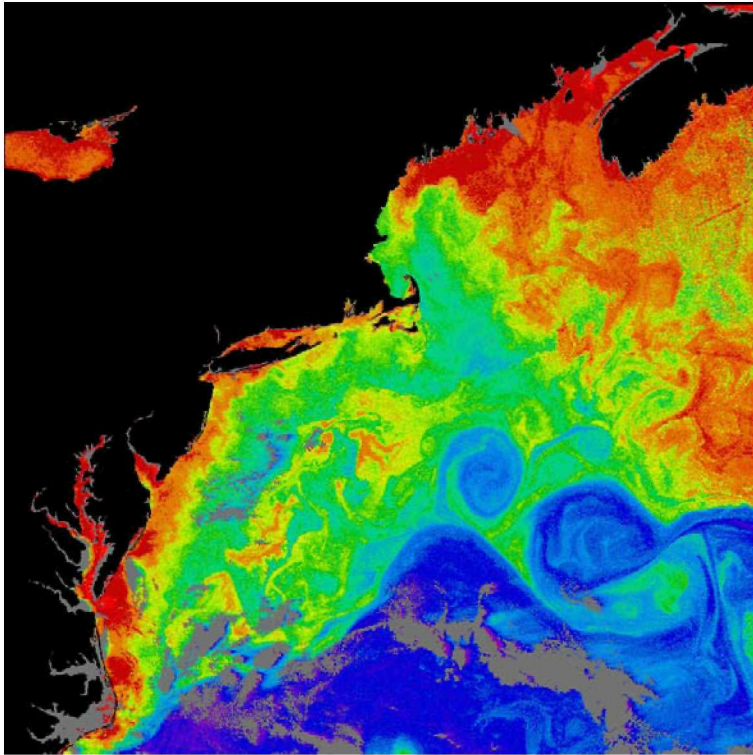
Pelegri & Csanady (1991)

c) possible trajectory over the double gyres



N pathway involves a figure of 8 :

5. Eddy/Fronts



Eddy Rectification

McGillicuddy & Robinson (1997)

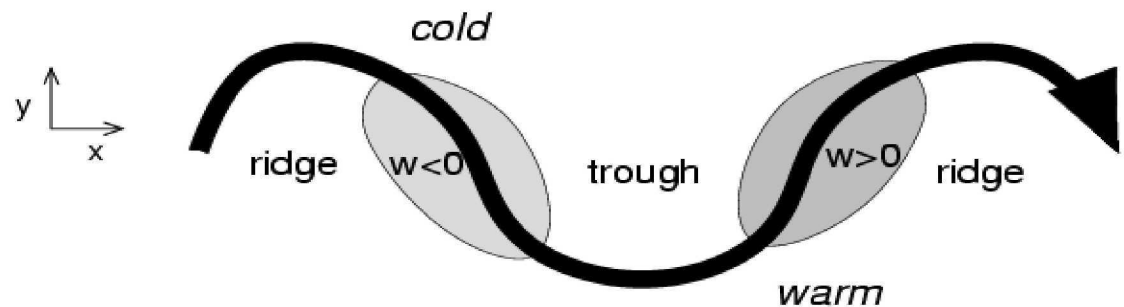
How does the process work?

Instability process

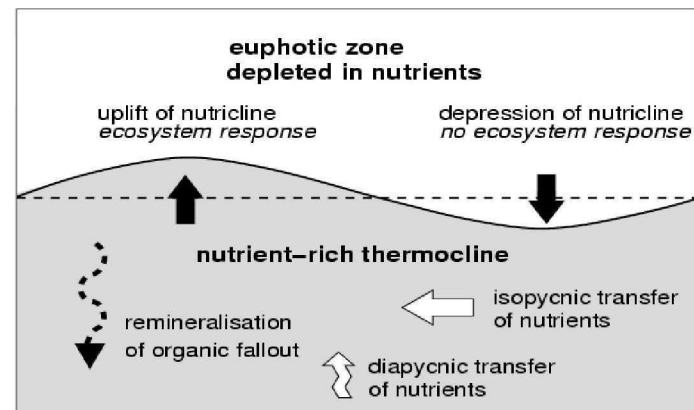
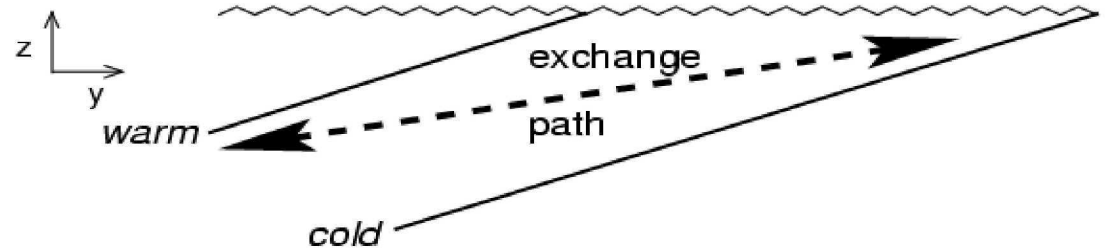
Warm waters *rise*
Cold waters *sink*

High N in surface waters of cold eddy is from lateral exchange, not local upwelling

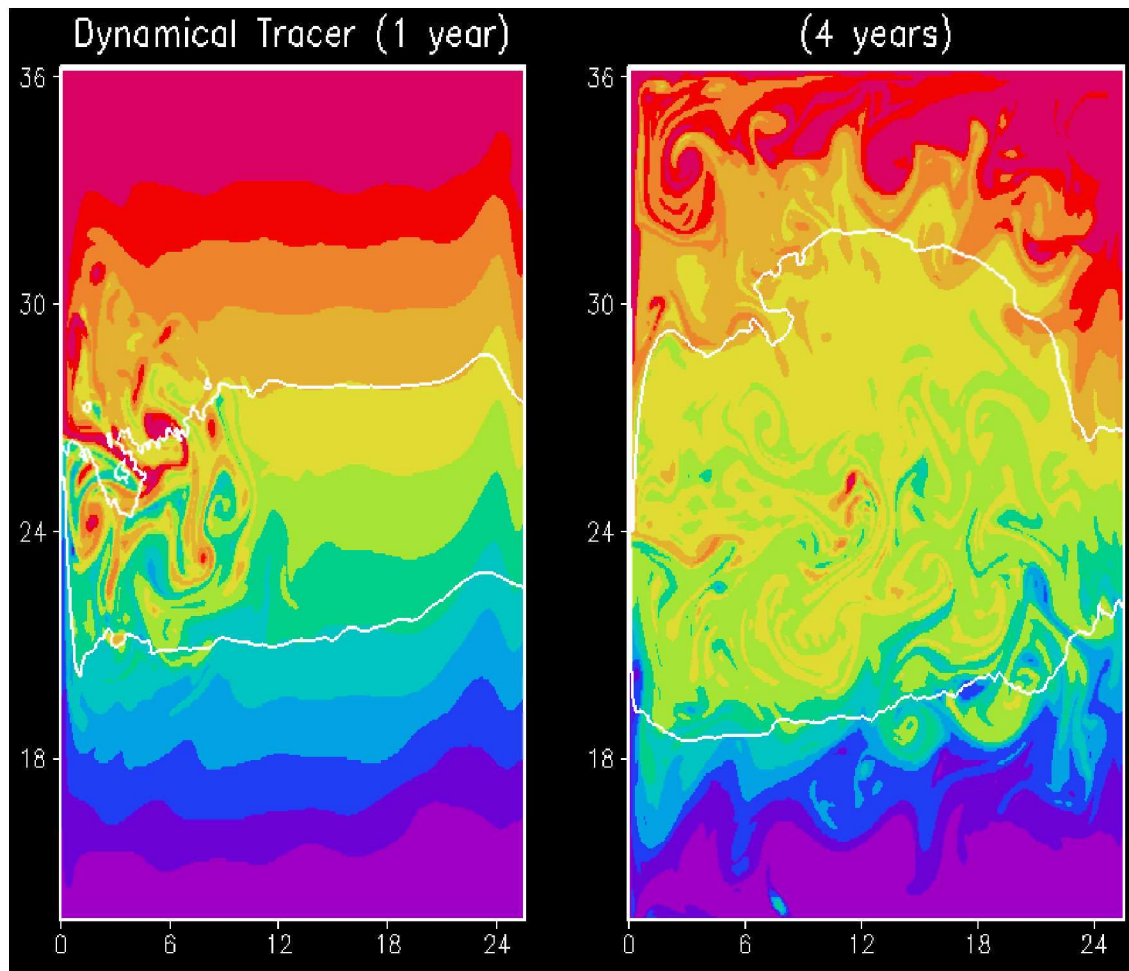
plan view of meandering jet



slantwise exchange of fluid



Idealised eddy example



Double wind-driven gyre

- 1/16 deg.

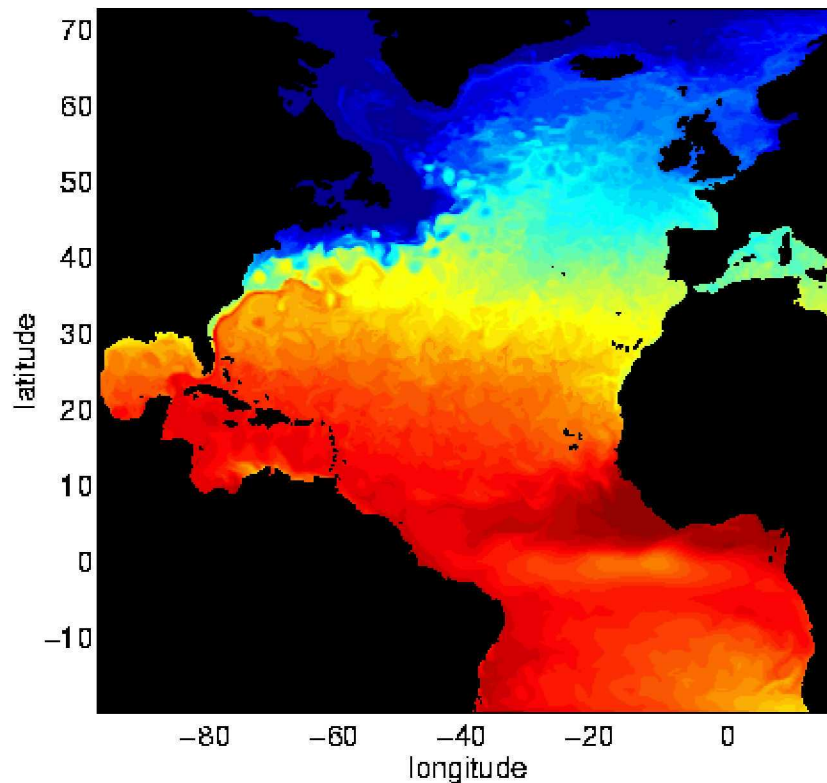
see dynamic tracer plotted on isopycnal

Wilson and Williams (2003)

Realistic GCM example

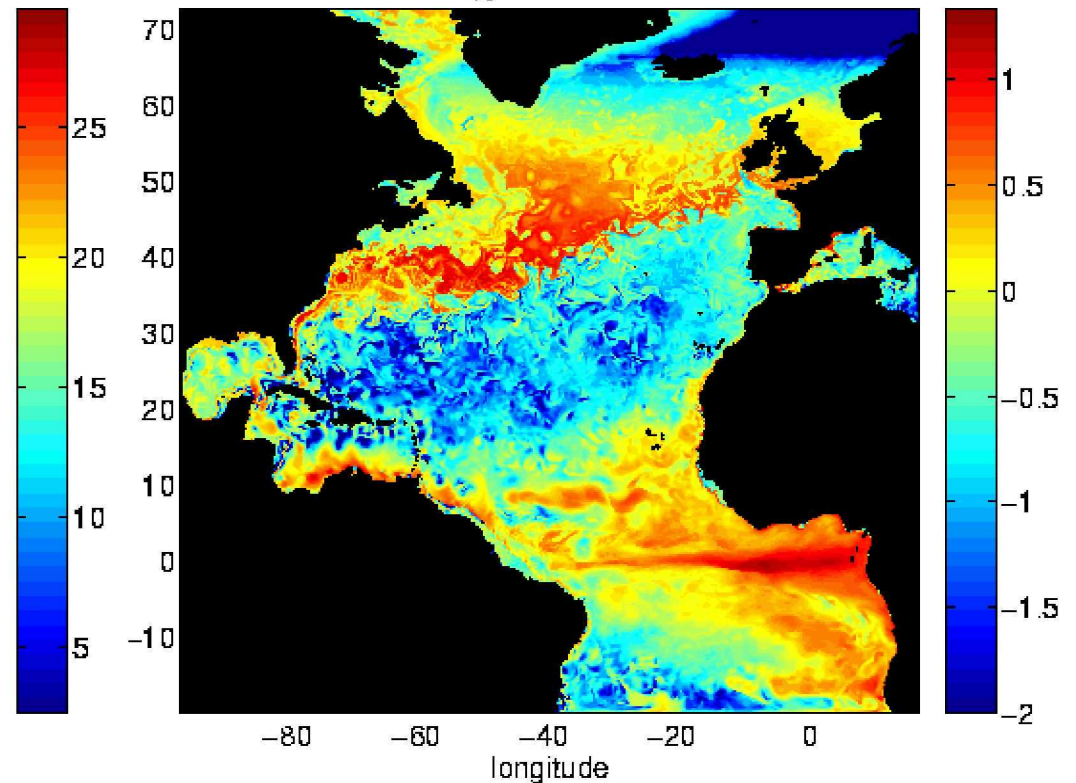
Surface T

Temperature (C) at 5 meters, 06 Jan 1993



New Production

New Production, \log_{10} (mmol N/m²/day), 06 Jan 1993



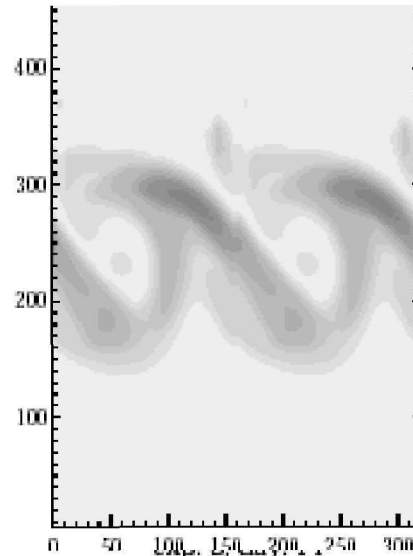
McGillicuddy et al. (2003)

1/10°, Los Alamos, 40 levels. 6 hr ECMWF winds 1985-1991, 11 year integration. N relaxed to climatology on 30 day timescale.

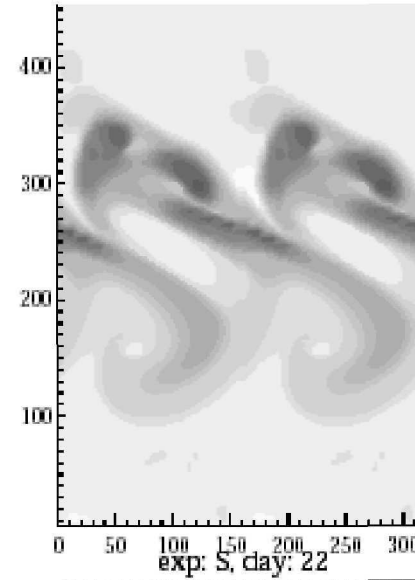
Fronts

1/3°

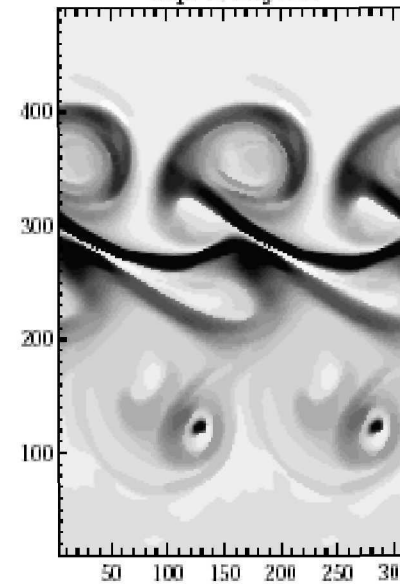
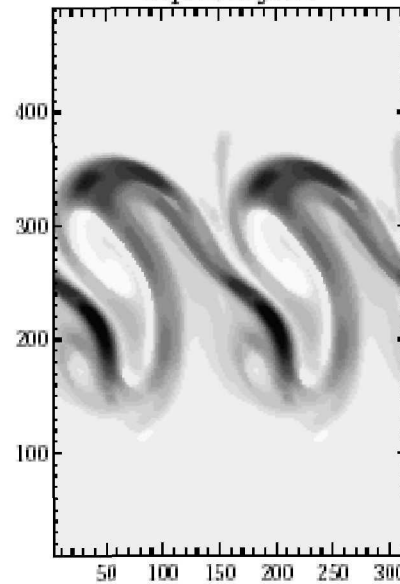
day 14



day 24



1/12°



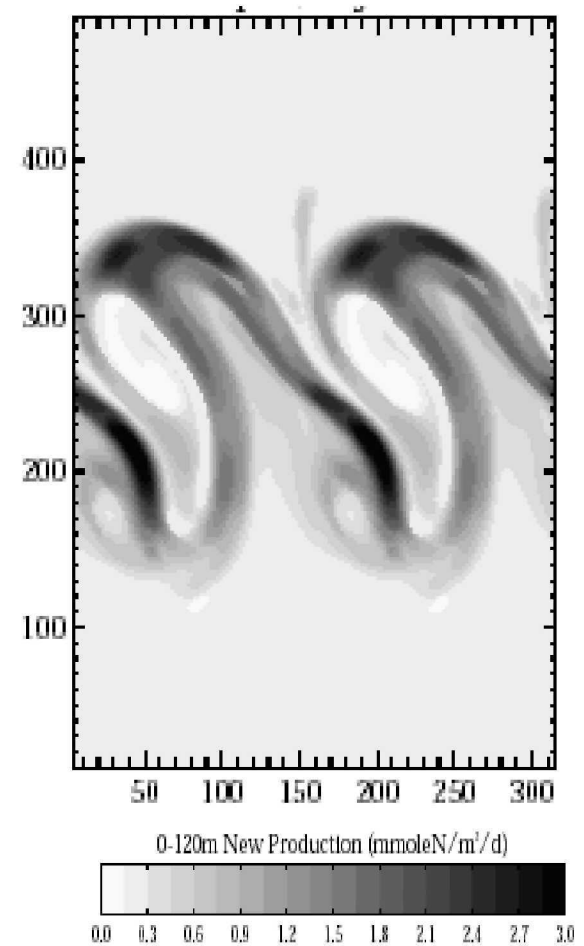
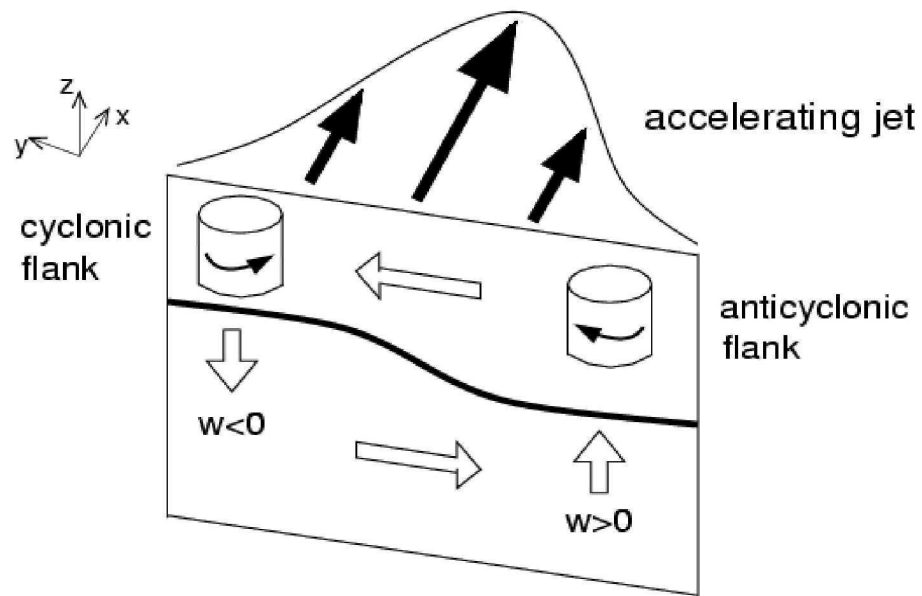
0-120m New Production (mmole N/m²/d)



Levy et al. (2001)

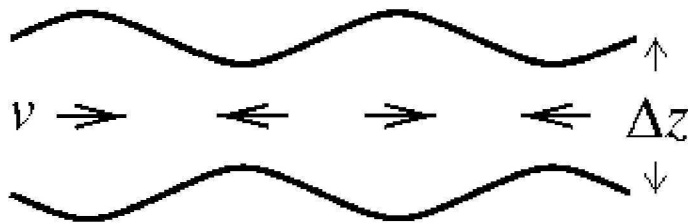
Frontal circulation

embedded frontal-scale circulation

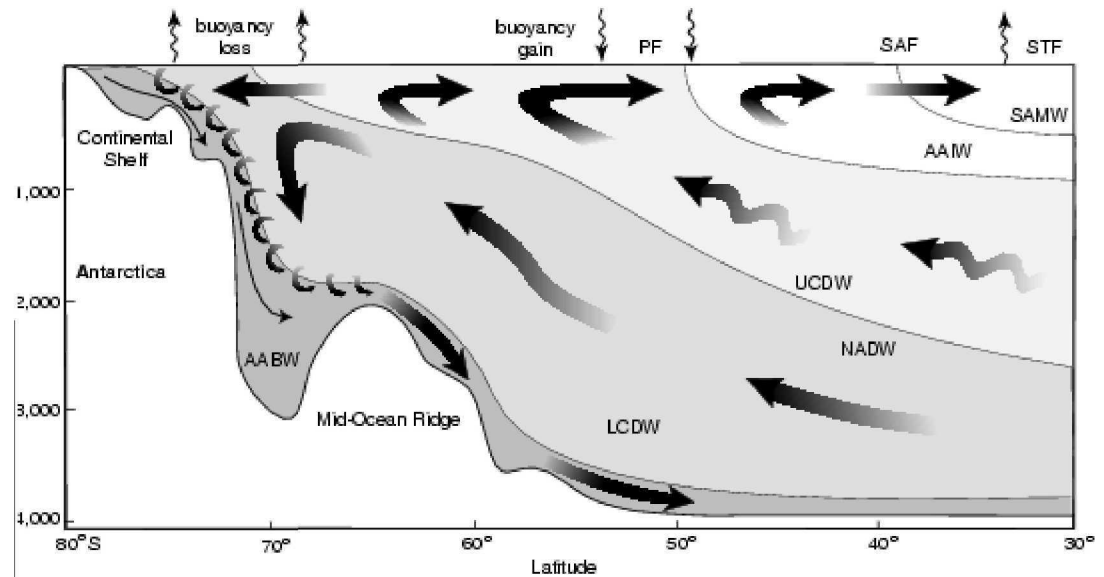
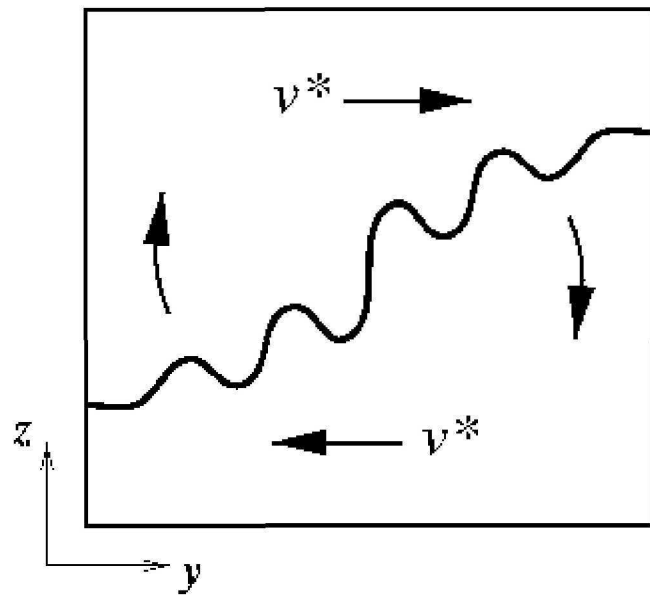


Eddy lateral transfer

a) eddy transport from velocity and thickness oscillations



b) eddy transport arising from slumping of isopycnals



Eddy transport probably crucial for southwards spreading in S. Ocean

6. Temporal variability

Steady state view omits
significant variability:

- Coupled atmosphere-ocean
- Atmosphere forced changes

convection

overturning

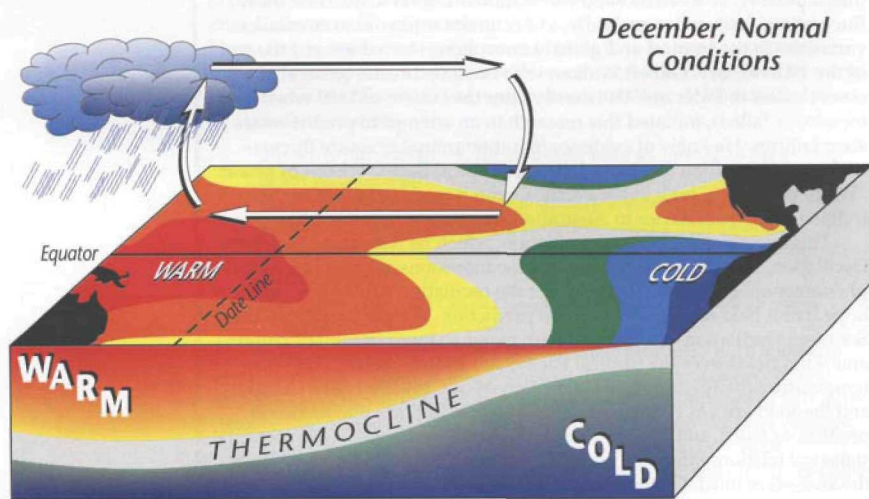
examples:

El Nino

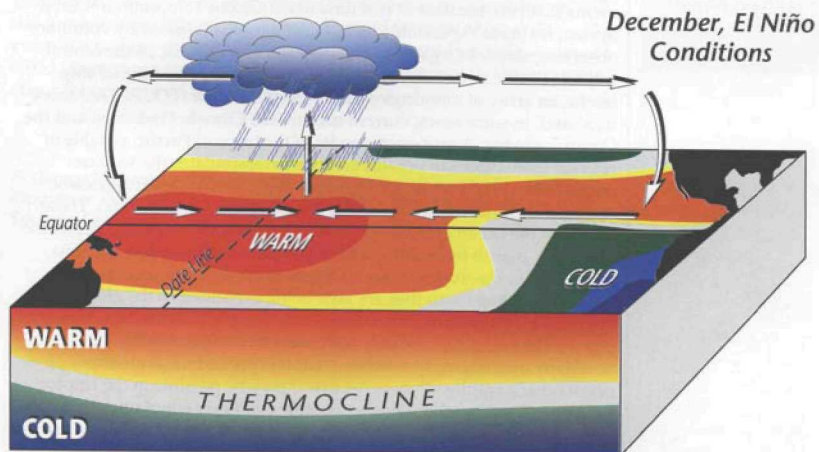
NAO

Mediterranean

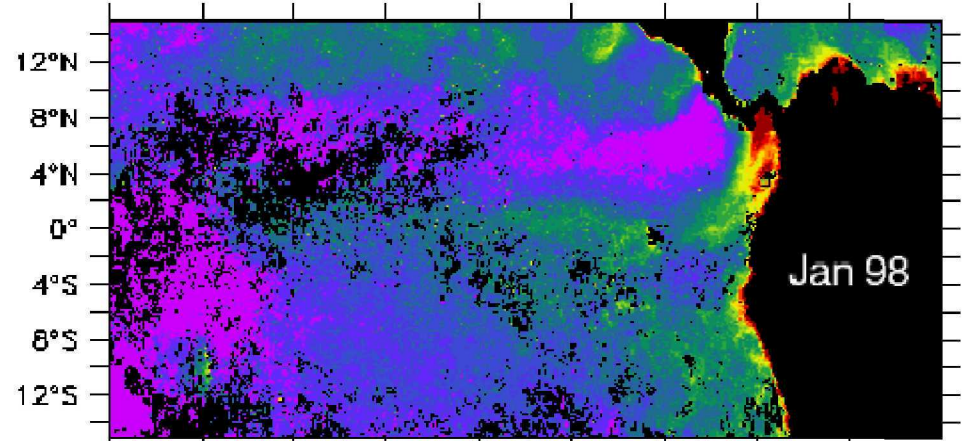
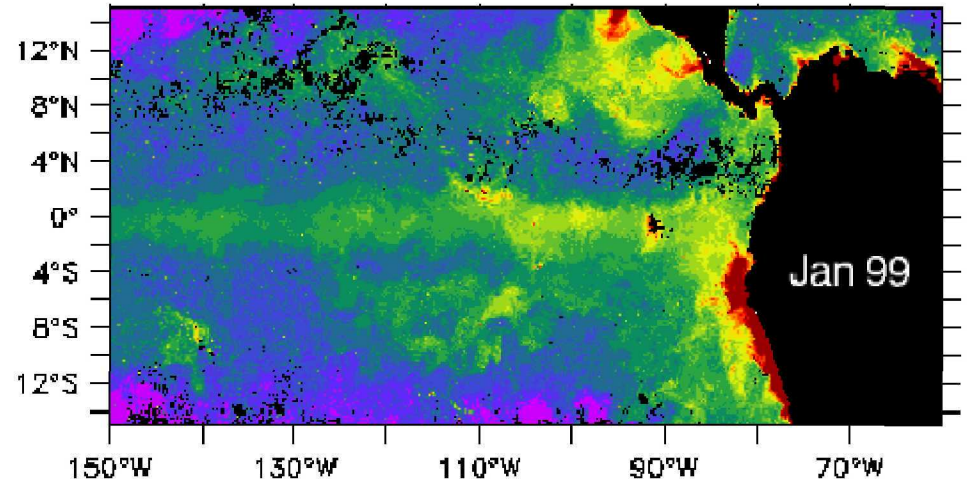
El Niño



Jayne Doucette/WHOI Graphics



Jayne Doucette/WHOI Graphics

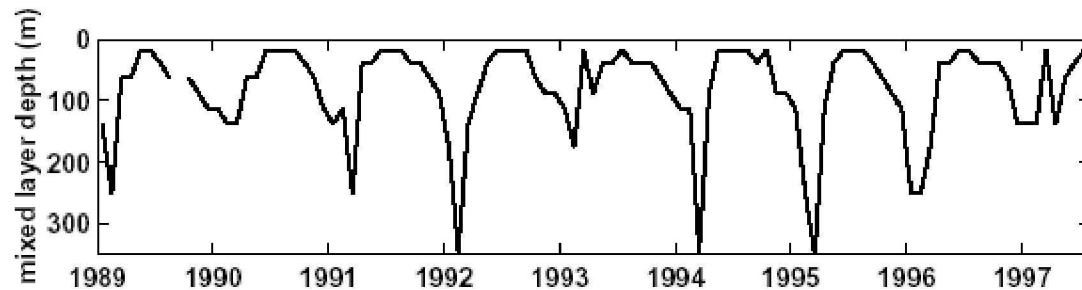


Philander (1990)

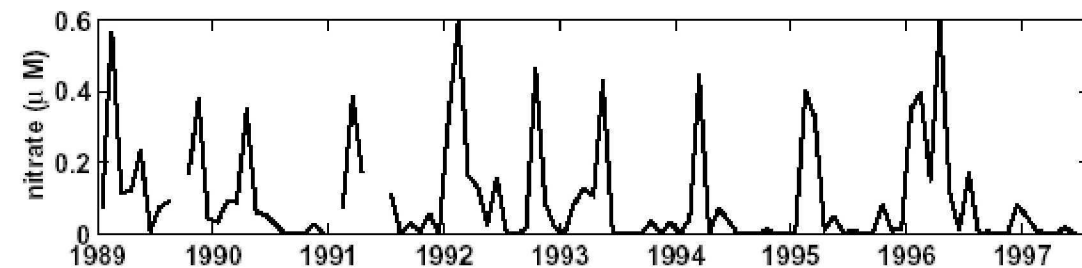
Chlorophyll a (mg/m³)

Bermuda Atlantic Time Series

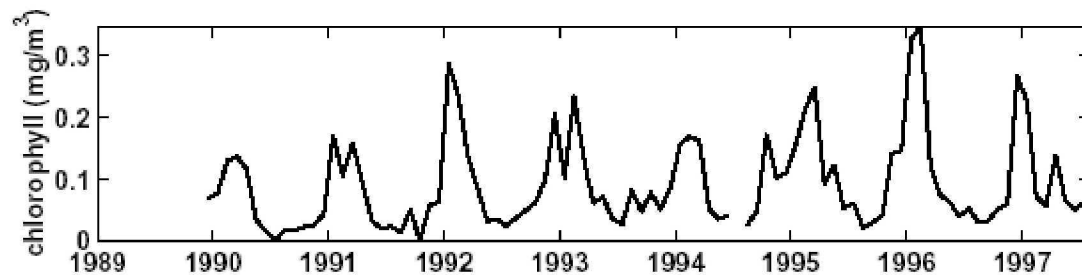
Mixed
layer
thickness



Nitrate

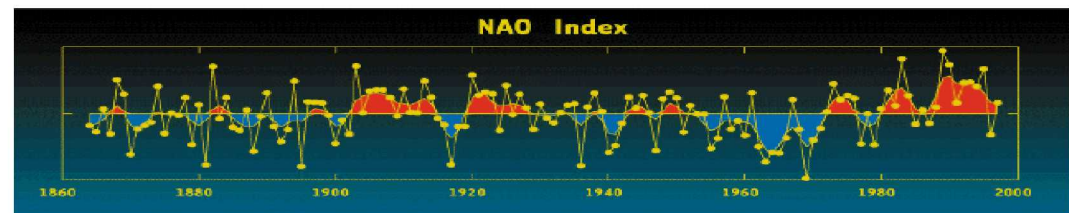
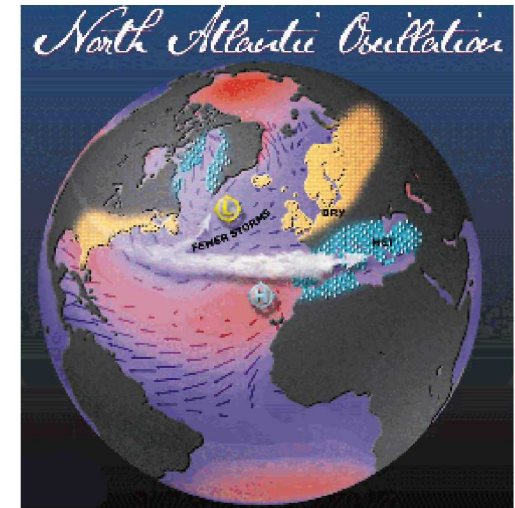
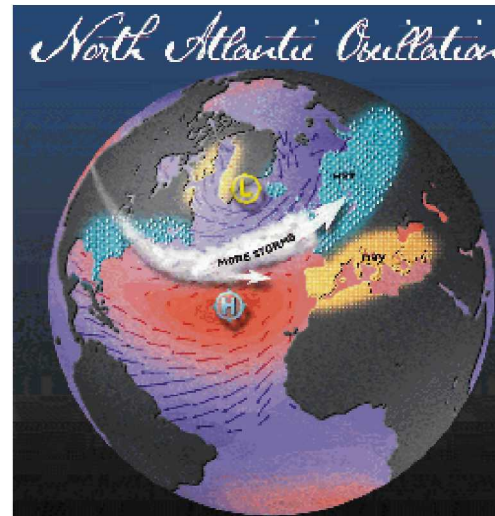


Chlorophyll



NAO

Dominant N. Atlantic atmospheric mode captures 30% of variability

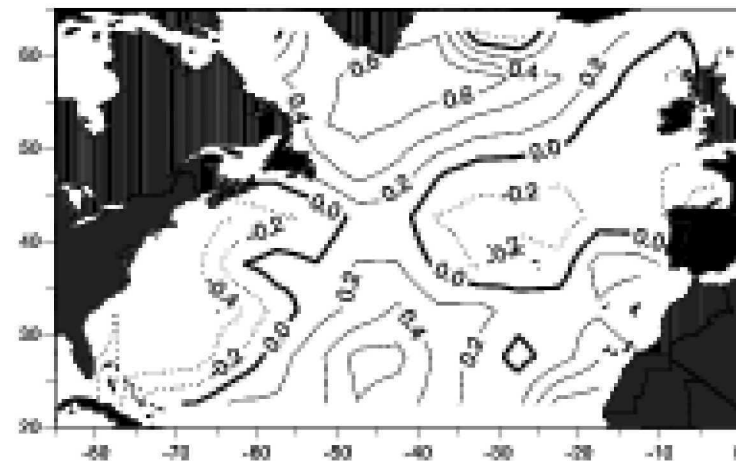


<http://www.ideo.columbia.edu/NAO/>

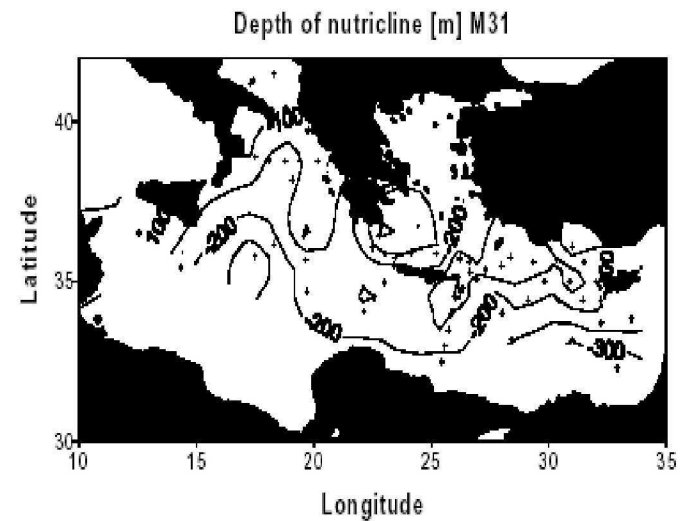
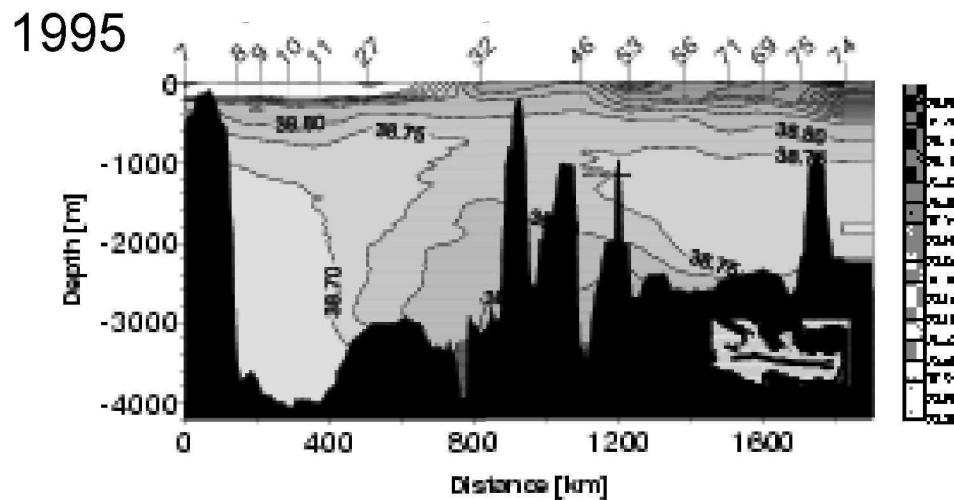
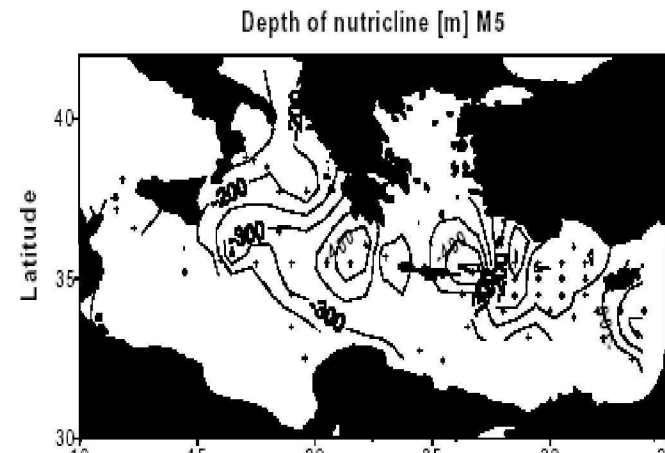
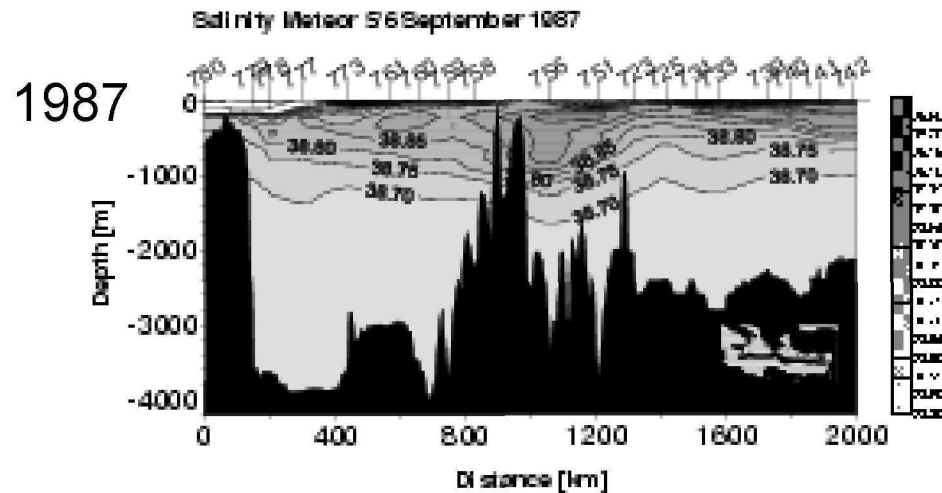
For nutrient supply, simple model predicts strong correlation over W. Atlantic, but not E. Atlantic

Williams et al. (2000)

d) Correlation: NAO and Convective nitrate supply

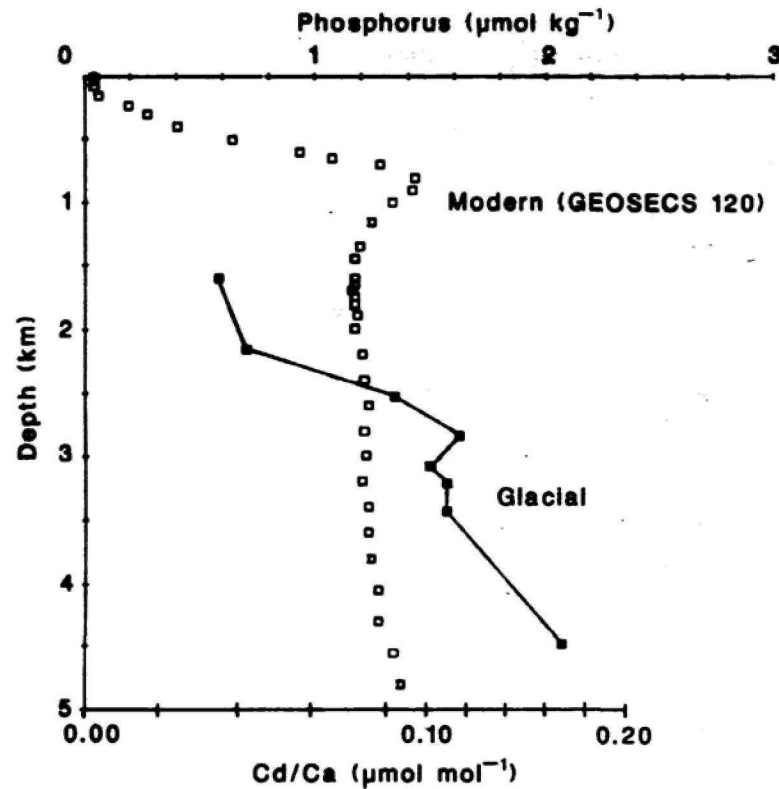


Mediterranean: abrupt change



Roether et al. (1996); Klein et al. (1999)

Glacial/ Interglacial in Atlantic



Glacial reconstruction based
on Cd/Ca ratios

Boyle & Keigwin (1987)

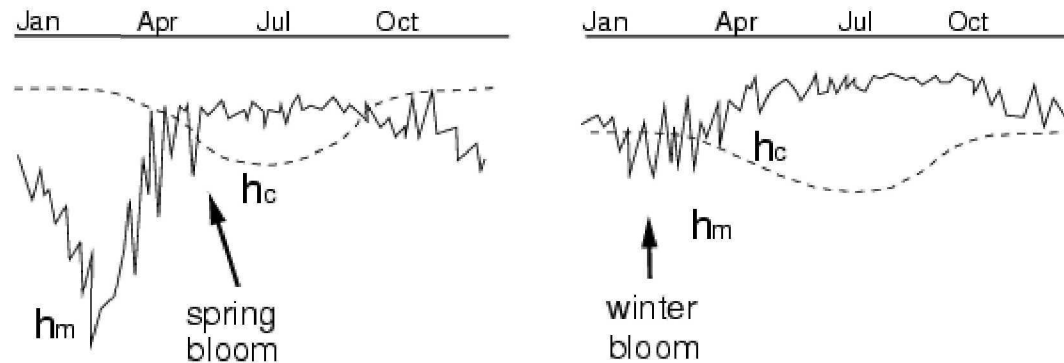
Conclusions

- Basin contrasts in N controlled by overturning circulation. Productive high latitudes due to convection & upwelling
- Separate gyres connected by Ekman flows and boundary currents
- Eddy & frontal productivity due to vertical & *lateral* transfers
- Significant interannual variability, usually linked to atmospheric forcing

Challenges:

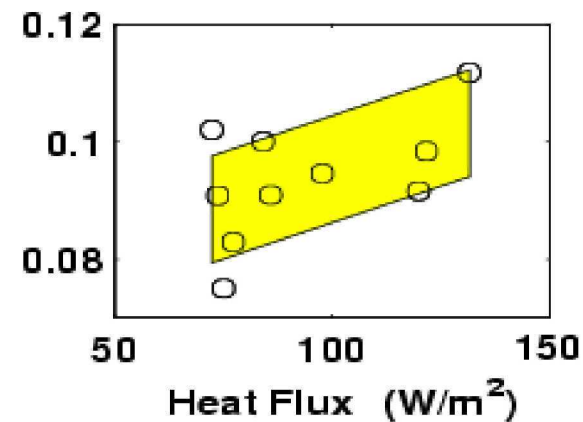
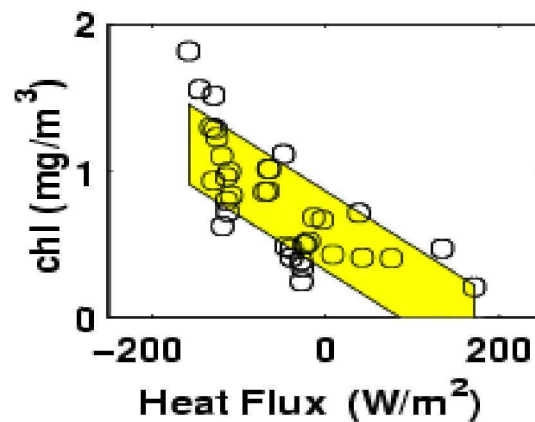
- acquire a more quantitative assessment
- separate perturbation effects & long term changes
- understand mechanistic links with the atmospheric

Convective variability



subpolar

subtropical



Follows & Dutkiewicz (2002)