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International Advanced School  
on  
Space Weather  
2-19 May 2006

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### *The Cosmic Ray Record*

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ITALY

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These lecture notes are intended only for distribution to participants



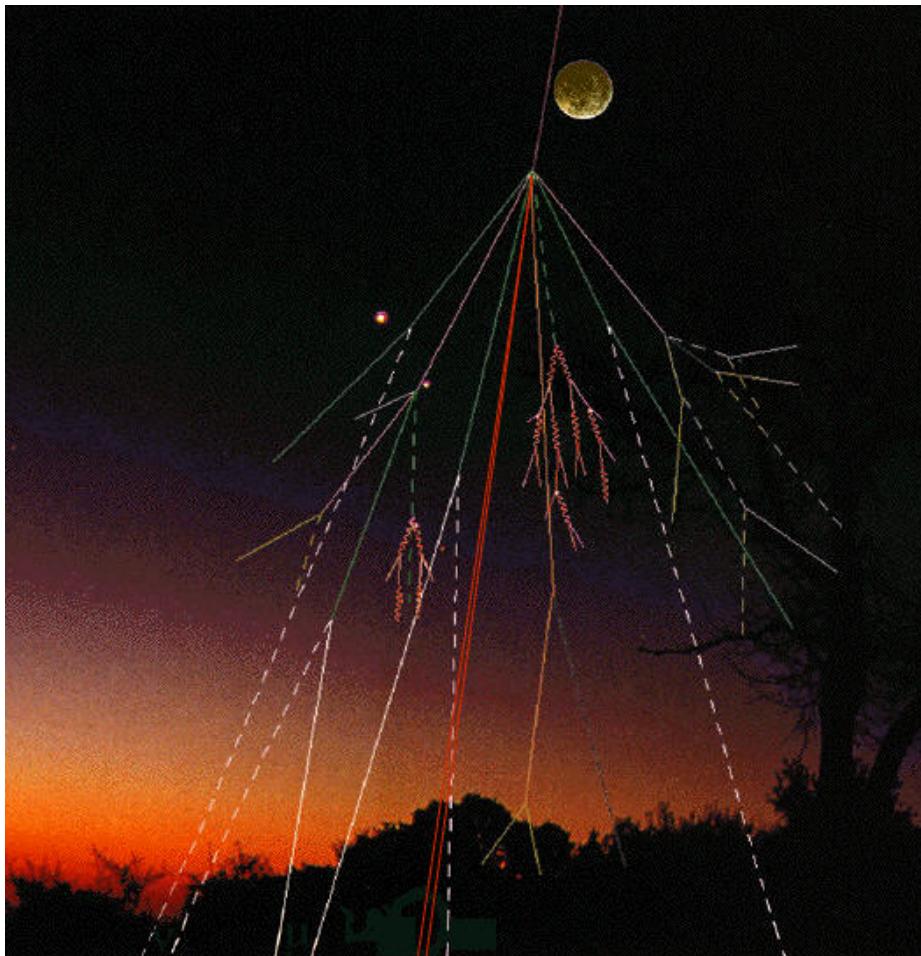
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# The cosmic ray record

Claudio Tuniz

The Abdus Salam International Centre for Theoretical Physics,  
Trieste, Italy

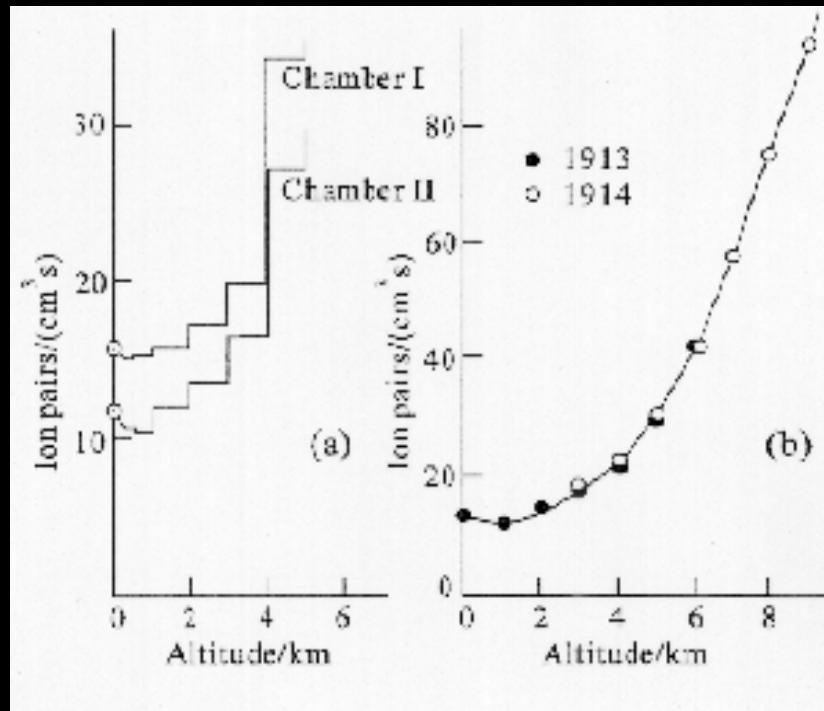
# A nuclear physicist's viewpoint



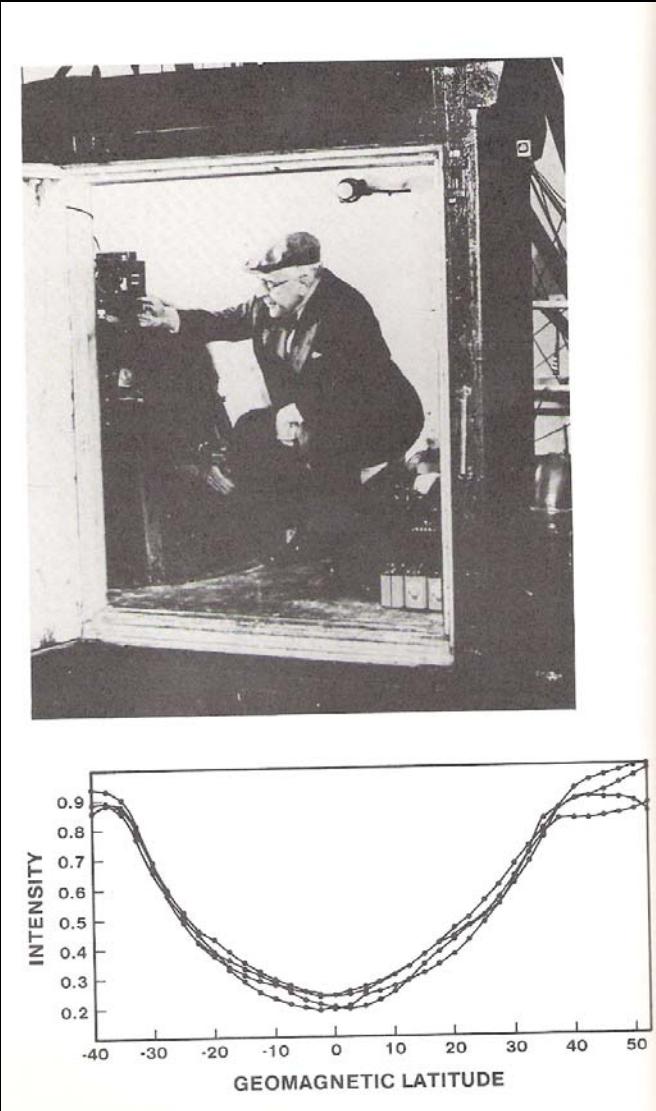
## Lecture outline:

- Cosmic rays (ion beam)
- Solar system matter (target)
- Nuclear reactions (radionuclide production)
- Counting rare radio-nuclides
- Extraterrestrial archives
- Terrestrial archives
- Solar variability

# Cosmic rays: history

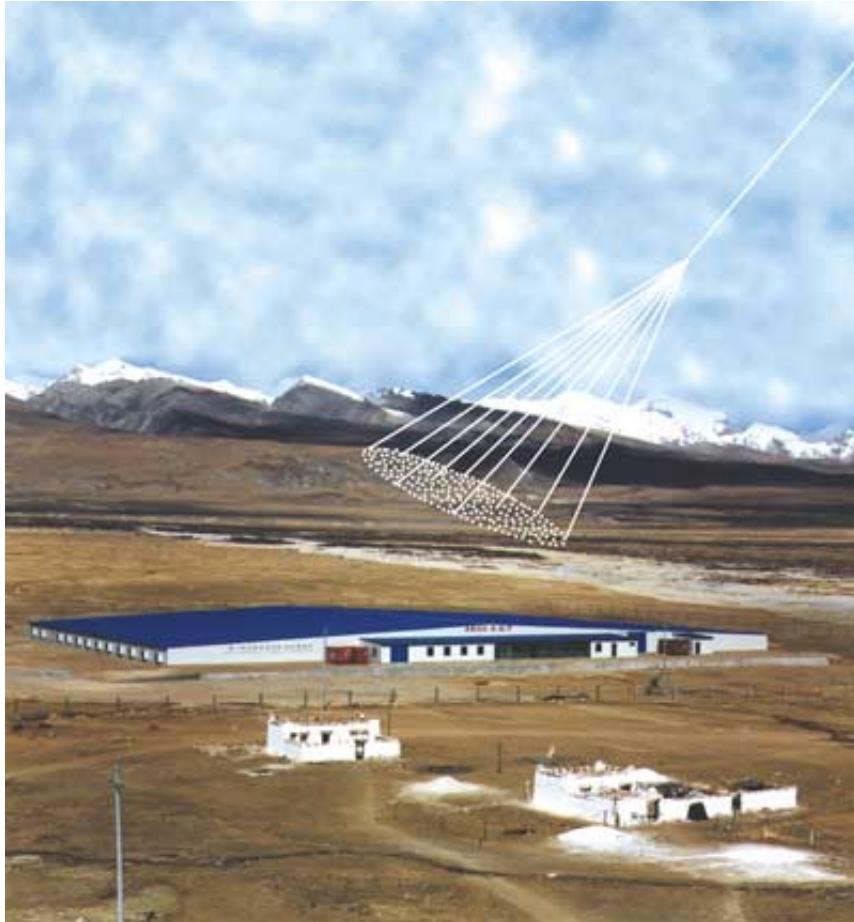


# Cosmic ray: history



From M W Friedlander, Cosmic Rays

# Last generation CR detectors



ARGO experiment in Tibet. A simulated particle shower strikes the ARGO detector  
18,500 detectors will cover an area as large as a soccer field.



NEMO experiment in Italy. Apparatus to detect the Cerenkov light produced by the passage of neutrinos through water to be installed undersea at a depth of two 5 kilometers .

# Cosmic rays: energy and flux



*Table 5.1* Comparison of Solar Wind, SCR, and GCR Fluxes

	Solar Wind	SCR	GCR
Nucleon energies	0.3 – 3 keV	~1 – 100 MeV	~0.1 – >10 GeV
Electron energies	~1 – 100 eV	<0.1 – 1 MeV	~0.1 – >10 GeV
Proton fluxes ( $\text{cm}^{-2} \text{ sec}^{-1}$ )	$\sim 3 \times 10^8$	100	2 – 4
Penetration depth in solid matter	$\mu\text{m}$	cm	m

*Source:* Vaniman, D. et al., in *Lunar Sourcebook: A User's Guide to the Moon*, Heiken, D.H. et al., Eds., Cambridge University Press, London, 1991, 27. With permission.

# Cosmic rays: elemental abundance



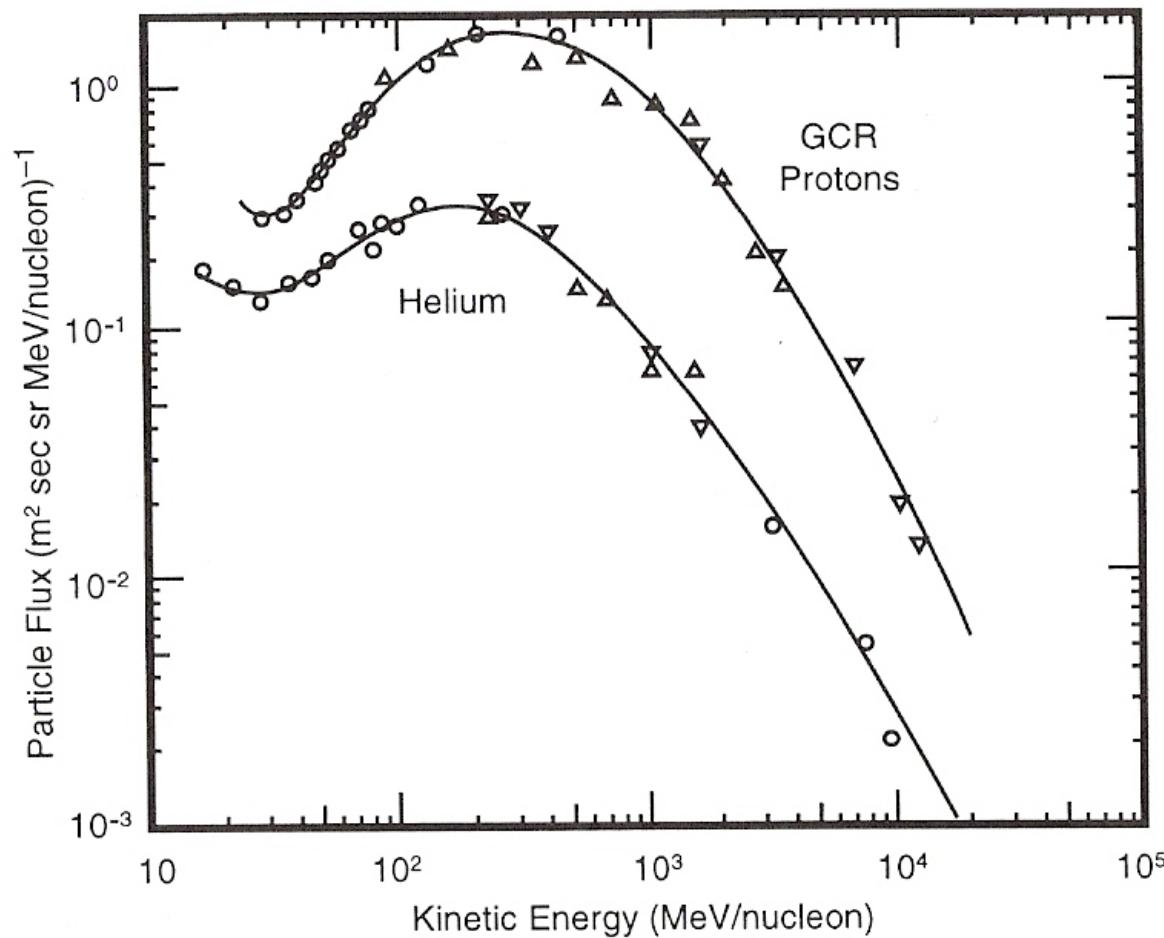
Accelerator mass Spectrometry  
C. Tuniz et al., CRC Press, 1998

**Table 5.2** Elemental Composition of GCR Compared with the Abundance of the Same Elements in the Universe

Element	Atomic Number	Abundance (%)	
		GCR	Universe
H	1	93	91
He	2	6.3	9.1
Li, Be, B	3 – 5	0.10	0.4
C, N, O, F	6 – 9	0.42	0.14
Ne-K	10 – 19	0.14	0.014
Ca-Zn	20 – 30	0.04	$2 \times 10^{-3}$
Ga-U	31 – 92	$2 \times 10^{-6}$	$10^{-6}$

Source: Pomerantz, M.A. and Duggal, S.P., *Rev. Geophys. Space Phys.*, 12, 243, 1974. With permission.

# Cosmic rays: energy spectrum



Accelerator mass Spectrometry  
C. Tuniz et al., CRC Press, 1998

Figure 5.2 The measured energy spectra of GCR proton and helium fluxes at 1 AU.

# Galactic modulation

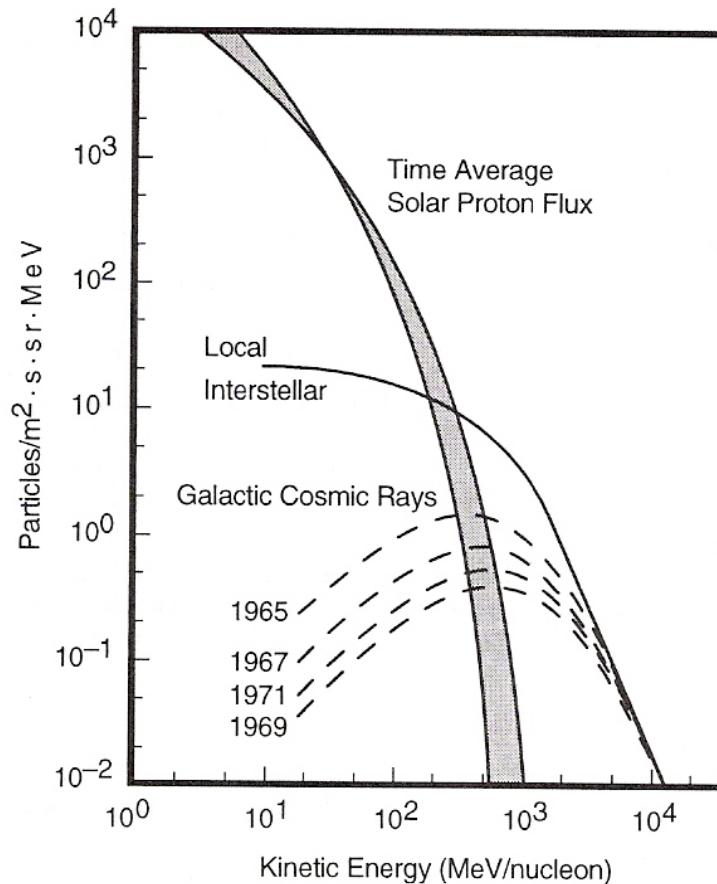


## 5.4.1 *Galactic modulation*

Galactic cosmic ray variability can result from several processes that are active outside the solar system on time scales on the order of  $10^6$  to  $10^9$  years:

- Fluctuations in the activity of the GCR sources
- Changes in acceleration and transport conditions in the interstellar medium during the period of  $10^6$  to  $10^7$  years that GCRs may spend in the galactic magnetic field
- Passage of the solar system through interstellar clouds
- Oscillation of the solar system through the galactic plane (Vanzani et al., 1987)

# Solar modulation



Accelerator mass Spectrometry  
C. Tuniz et al., CRC Press, 1998

**Figure 5.4** Calculated GCR spectra as a function of particle energy showing the effects of solar modulation. For comparison, the calculated GCR spectrum in interstellar space and the time averaged SCR proton flux are also shown. (Adapted from Castagnoli, G. and Lal, D., *Radiocarbon*, 22, 133, 1980; Reedy, R.C. et al., *Ann. Rev. Nucl. Part. Sci.*, 33, 505, 1983.)

# Cosmic ray interactions

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- Nuclear reactions [light GCR & SCR]
  - Radioactive nuclides [ $^{10}\text{Be}$ ,  $^{26}\text{Al}$ ,  $^{53}\text{Mn}$ , ...]
  - Stable nuclides [noble gases]
- Ionization energy loss [SCR, heavy GCR]
  - Nuclear tracks
  - Thermo luminescence
- Implantation [solar wind]

# Nuclear reactions by cosmic rays

## (radionuclide production)



### 7.1 Atmospheric production of cosmogenic nuclides

The radionuclide production rate,  $P(h,\theta,t)$ , at altitude,  $h$ , and geomagnetic latitude,  $\theta$ , at time  $t$  is given by the general formula:

$$\text{Production rate} \quad P(h,N,t) = \int_E \left\{ \sum_k \sum_j \theta_k(h,\theta,t,E) \sigma_{jk}(E) N_j \right\} dE \quad (7.1)$$

where:

$\theta_k$  = differential energy spectrum of component  $k$  of the cosmic ray flux

$N_j$  = atom abundance of target nuclide  $j$  in the irradiated material

$\sigma_{jk}$  = cross-section for the production of the radionuclide in the collision of particle  $k$  of energy  $E$  with target nuclide  $j$

# Terrestrial production



Accelerator mass Spectrometry  
C. Tuniz et al., CRC Press, 1998

**Table 7.1** Cosmogenic Radionuclides  
Produced in Terrestrial Matter

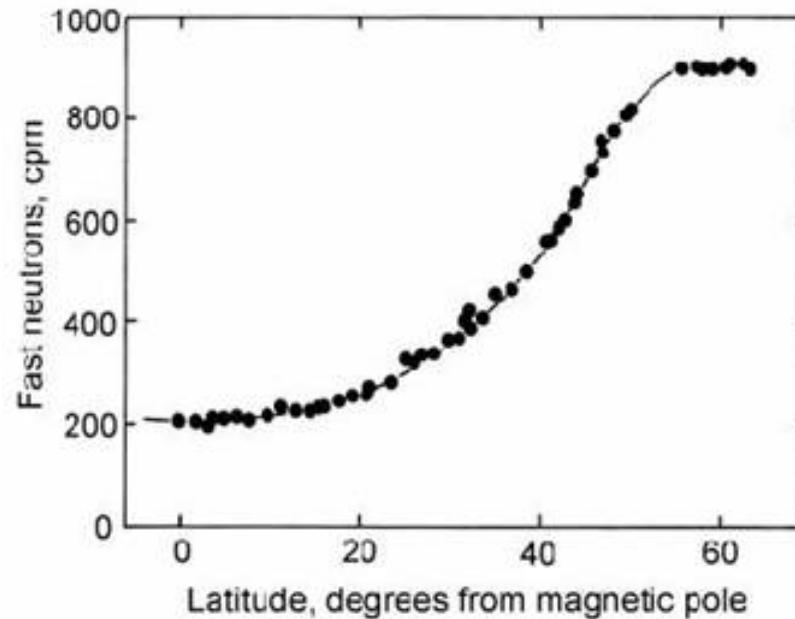
Radionuclide	Half-Life ( $t_{1/2}$ )	Main Terrestrial Target Elements	
		Atmosphere	Lithosphere
$^3\text{H}$	12.3 a	N, O	O, Mg, Si, Fe
$^7\text{Be}$	53.3 d	N, O	O, Mg, si, Fe
$^{10}\text{Be}$	1.51 Ma	N, O	O, Si, Mg, Fe
$^{14}\text{C}$	5730 a	N, O	O, Mg, Si, Fe
$^{22}\text{Na}$	2.61 a	Ar	Mg, Al, Si, Fe
$^{26}\text{Al}$	720 ka	Ar	Cl, K, Ca, Fe
$^{32}\text{Si}$	~140 a	Ar	Cl, K, Ca, Fe
$^{36}\text{Cl}$	301 ka	Ar	Cl, K, Ca, Fe
$^{39}\text{Ar}$	269 a	Ar	K, Ca, Fe
$^{41}\text{Ca}$	103 ka	Kr	Ca, Ti, Fe
$^{81}\text{Kr}$	210 ka	Kr	Rb, Sr, Zr
$^{90}\text{Sr}$	28.8 a	Fission	
$^{129}\text{I}$	15.7 Ma	Xe	Te, Ba, Ce, La

*Note:* The listed isotopes have been measured by AMS or have been proposed to be measured by AMS.

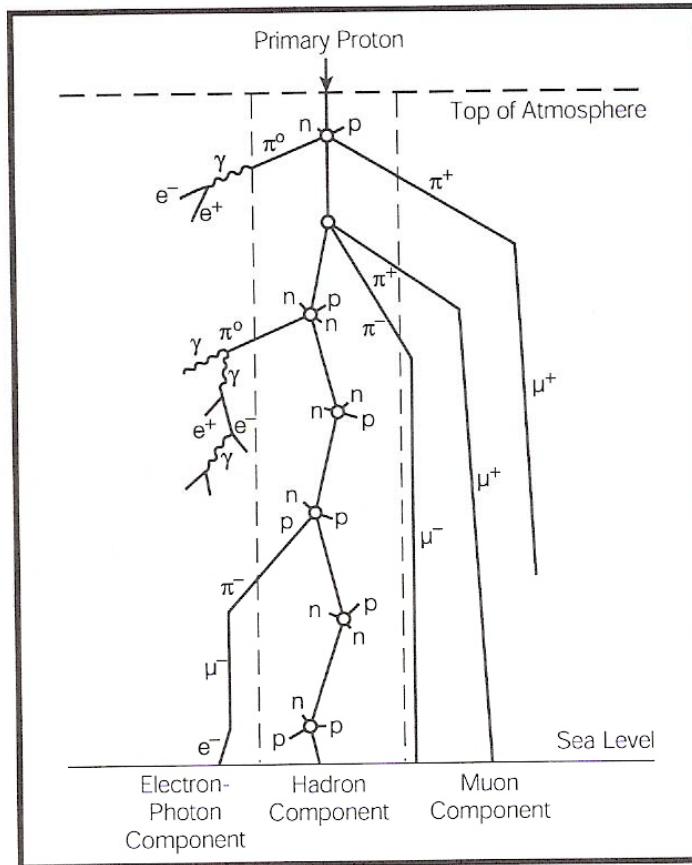
# Cosmogenic radionuclides

The rate of production of cosmogenic radionuclides depends on:

- 1) **latitude**
- 2) **geomagnetic field strength**
- 3) **solar activity**



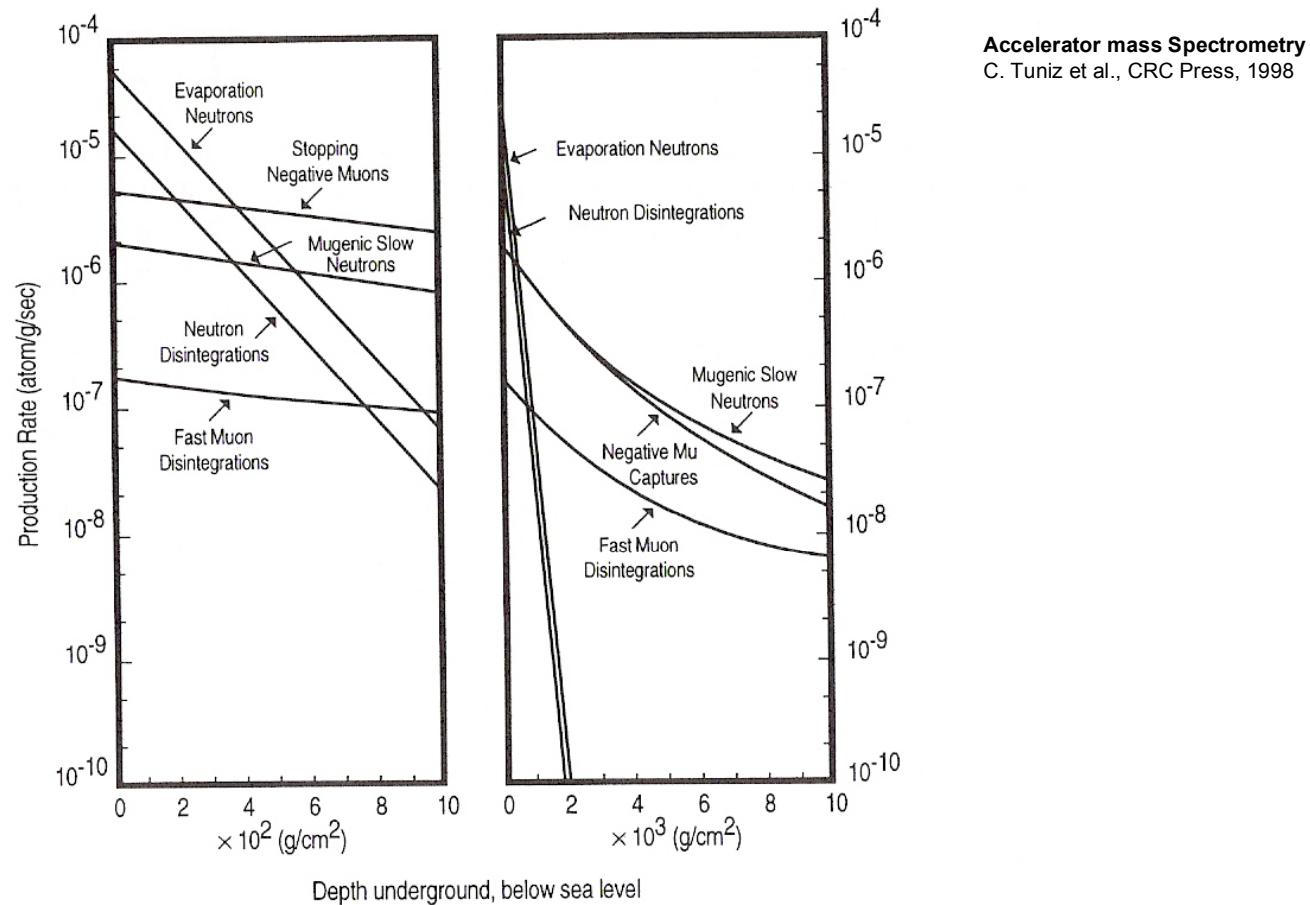
# Secondary cosmic rays



Accelerator mass Spectrometry  
C. Tuniz et al., CRC Press, 1998

**Figure 5.1** Schematic illustration of the development of the secondary cosmic ray flux produced in the atmosphere and the resultant flux at the Earth's surface as a result of galactic cosmic ray bombardment. The three panels show the different particles generated in the electron, hadron (protons and neutrons), and muon components.

# Secondary CR reactions in the lithosphere

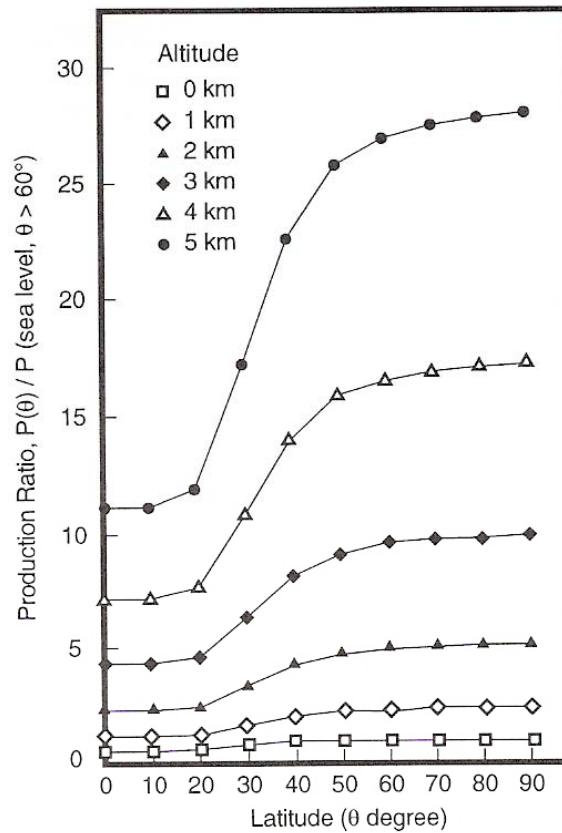


**Figure 7.2** Nuclear interaction rates due to nucleons and muons in the upper layers of the Earth's crust as a function of depth below the surface. The left-hand figure shows the production rate variations on an expanded scale for depths up to about 3 m. (From Lal, D., *Earth Planet. Sci. Lett.*, 104, 424, 1991. With permission.)

# Production of cosmogenic isotopes in the lithosphere



Accelerator mass Spectrometry  
C. Tuniz et al., CRC Press, 1998

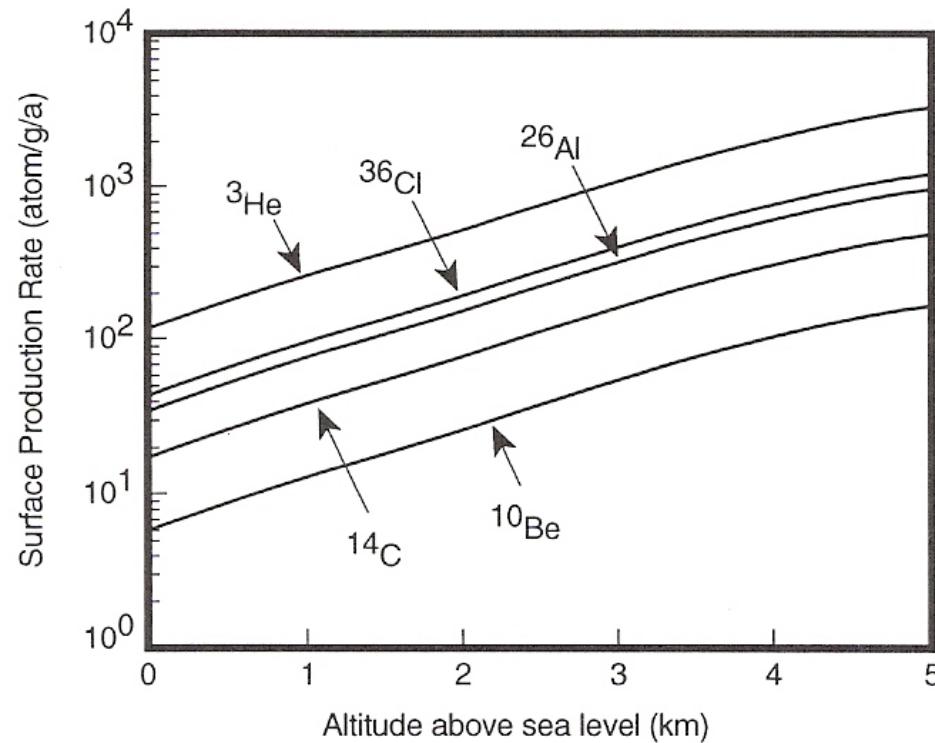


**Figure 11.1** Variation of *in situ* production rates as a function of latitude and altitude,  $P(\theta)$ , normalized to the sea-level, high-latitude production rate ( $P$ ) (latitude  $> 60^\circ$ ). The polynomial parameterization given by Lal (1991) for latitude and altitude variations of the cosmic ray flux were used (see Figure 7.1). (Courtesy of J. Klein, University of Pennsylvania.)

# Production of cosmogenic isotopes in the lithosphere



Accelerator mass Spectrometry  
C. Tuniz et al., CRC Press, 1998



**Figure 7.7** Calculated terrestrial production rates of *in situ* cosmogenic nuclides in surface rocks as a function of altitude based on numerical simulation computer codes using an incident GCR proton flux of 4.6 protons/cm<sup>2</sup>/sec for energies >10 MeV. The target for  $^{36}\text{Cl}$  is CaO and SiO<sub>2</sub> for the others. (From Masarik, J. and Reedy, R.C., *Earth Planet. Sci. Lett.*, 136, 381, 1994. With permission.)

# Measured and predicted depth dependence for $^{10}\text{Be}$ and $^{36}\text{Cl}$ (Elmore et al, Purdue University)

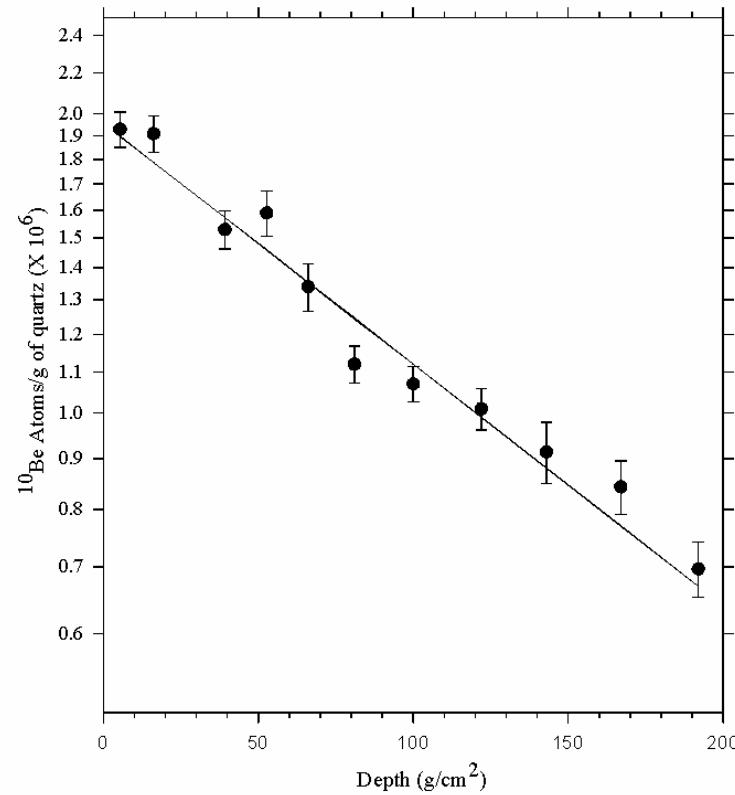


Figure 2

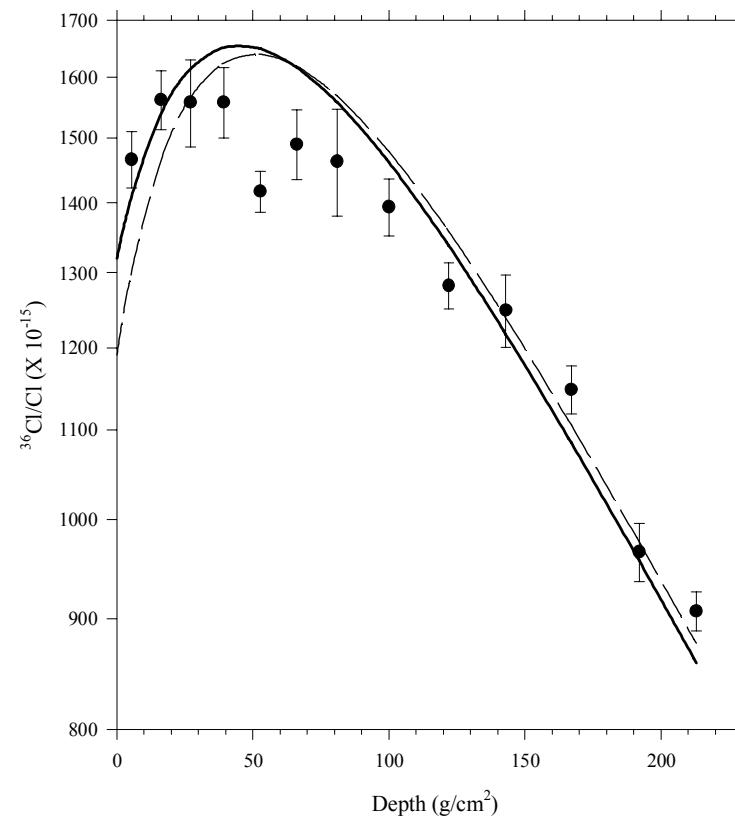
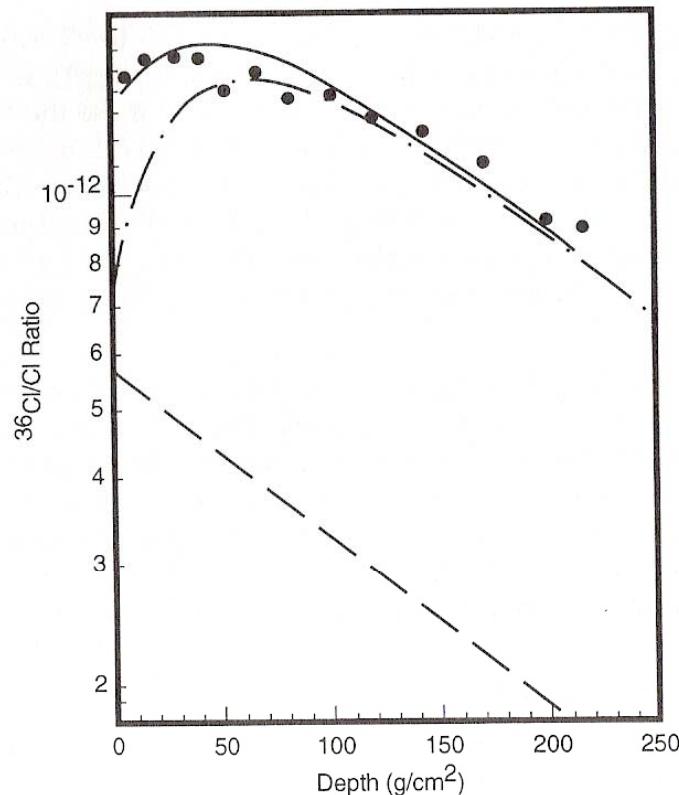


Figure 1

# $^{36}\text{Cl}$ production in a boulder



Accelerator mass Spectrometry  
C. Tuniz et al., CRC Press, 1998



**Figure 7.6** Measured and predicted depth dependence of  $^{36}\text{Cl}/\text{Cl}$  in a 90-cm core taken from a granodiorite erratic boulder at Bishop Creek, California. The solid line is a model calculation normalized to the data points. The dot-dash line is the predicted ratio from production of  $^{36}\text{Cl}$  from neutron capture on Cl, and the dashed line is the calculated depth dependence for  $^{36}\text{Cl}$  production from spallation on Ca and K. The neutron capture component shows the distinct sub-surface maximum for production from neutrons, and the spallation component results from the exponential reduction in the flux of energetic secondary particles. (From Dep, L., Cosmogenic Radionuclide Production in Terrestrial Rocks: AMS Measurements and Monte Carlo Simulations, Ph.D. thesis, Department of Physics, Purdue University, W. Lafayette, IN, 1995.)

# Samples

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- Meteorites (asteroids, moon, mars)
- Cosmic dust (stratosphere)
- Cosmic spherules (ocean)
- Lunar samples
- Comet return mission
- Asteroid return mission
- Spacecraft components
  - Surveyor III (Apollo 12)

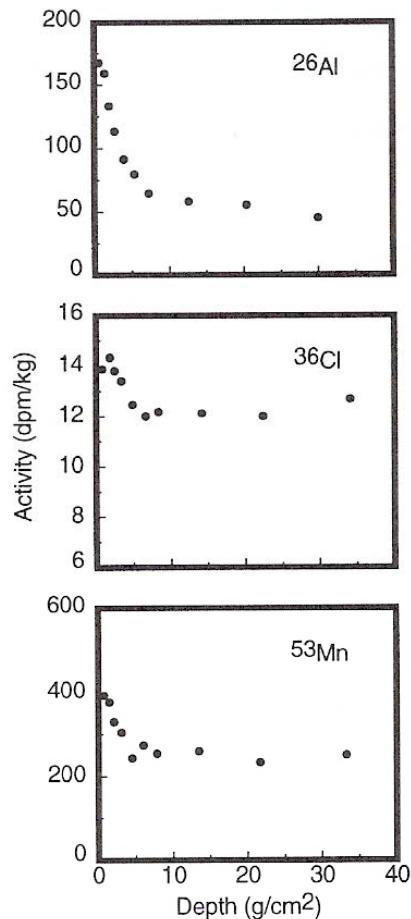
# Cosmic ray nuclear reactions: extraterrestrial material



*Production rate*

$$P_x(R,d) = \int_E \{ \sum_k \sum_j \phi_k(E,R,d) \sigma_{j,k}(E) N_j \} dE \quad (6.3)$$

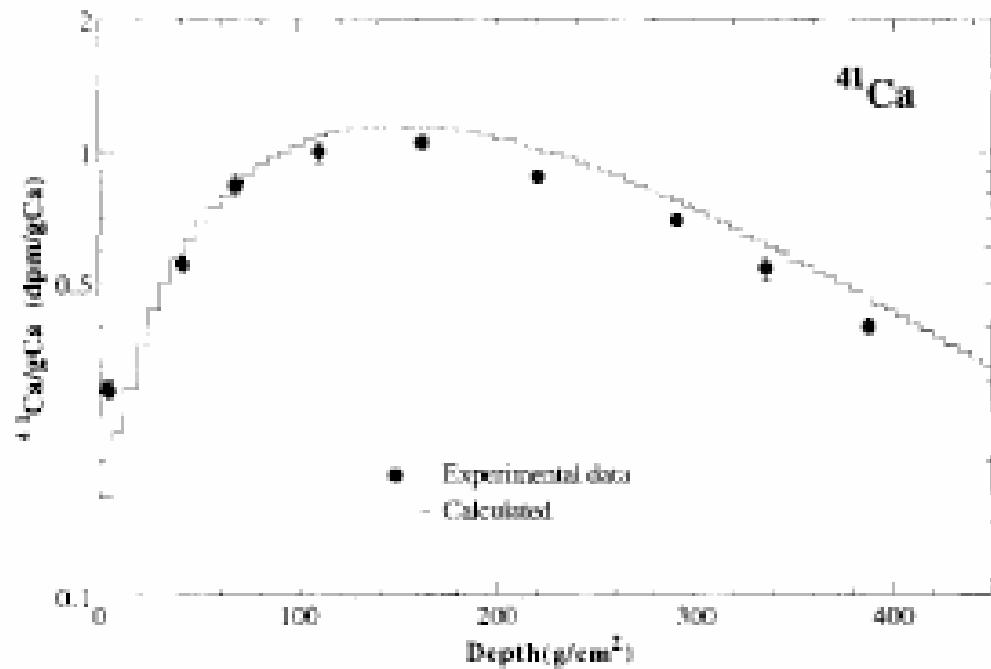
# Cosmogenic isotopes on the Moon



Accelerator mass Spectrometry  
C. Tuniz et al., CRC Press, 1998

**Figure 6.7** Surface profiles of  $^{26}\text{Al}$ ,  $^{36}\text{Cl}$ , and  $^{53}\text{Mn}$  in selected lunar cores and rocks. Depth range is from surface to about 15 cm. Activities increase sharply towards the surface, reflecting production by low-energy solar cosmic rays. (Profile of  $^{26}\text{Al}$  from Fruchter, J.C. et al., *Proc. 13th Lunar Planet. Sci. Conf.*, 13, 243, 1982;  $^{36}\text{Cl}$  from Nishiizumi, K. et al., *Proc. 19th Lunar Planet. Sci. Conf.*, XIX, 305, 1988;  $^{53}\text{Mn}$  from Nishiizumi, K. et al., *J. Geophys. Res.*, 88, 211, 1983. With permission.)

# Cosmogenic isotopes on the Moon



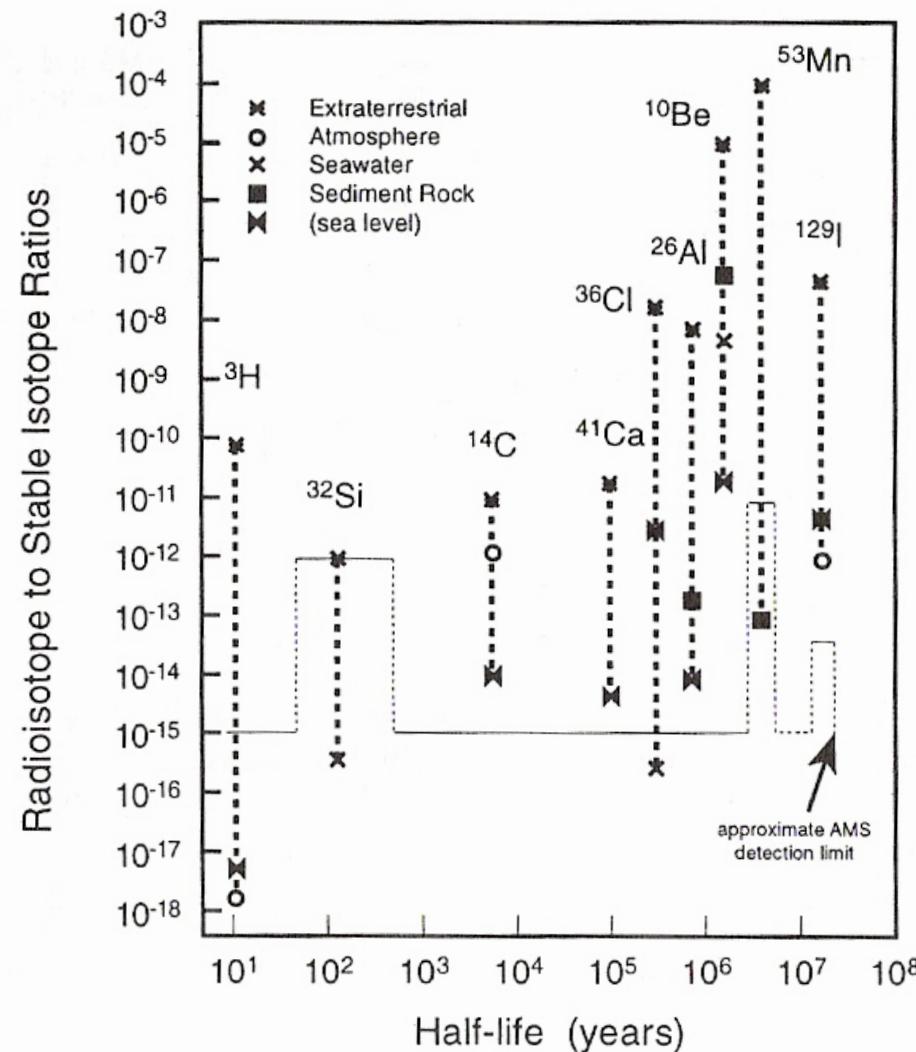
Nishiizumi et al, 1997.

Fig. 2. Low-energy neutron produced  $^{41}\text{Ca}$  activities in the Apollo 15 drill core. The experimental data have been normalized to Ca concentrations after subtraction of small amounts of spallation contribution (see text). The solid line is the calculated production rate using the LAHET Code System, assuming  $^{41}\text{Ca}$  from  $^{40}\text{Ca}$  ( $n,\gamma$ ) reaction.

# Long-lived cosmogenic radionuclides in solar system matter



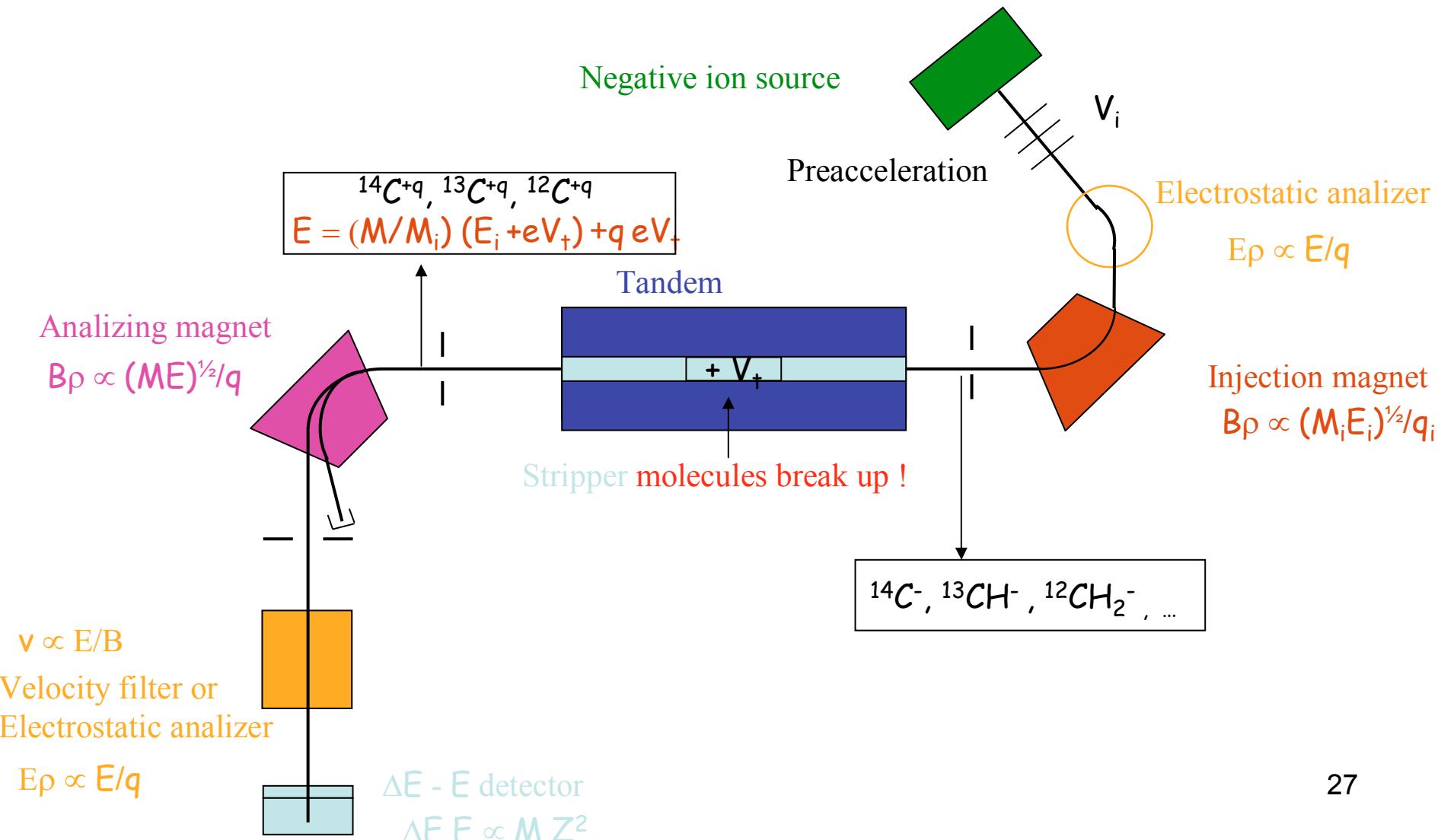
Accelerator mass Spectrometry  
C. Tuniz et al., CRC Press, 1998



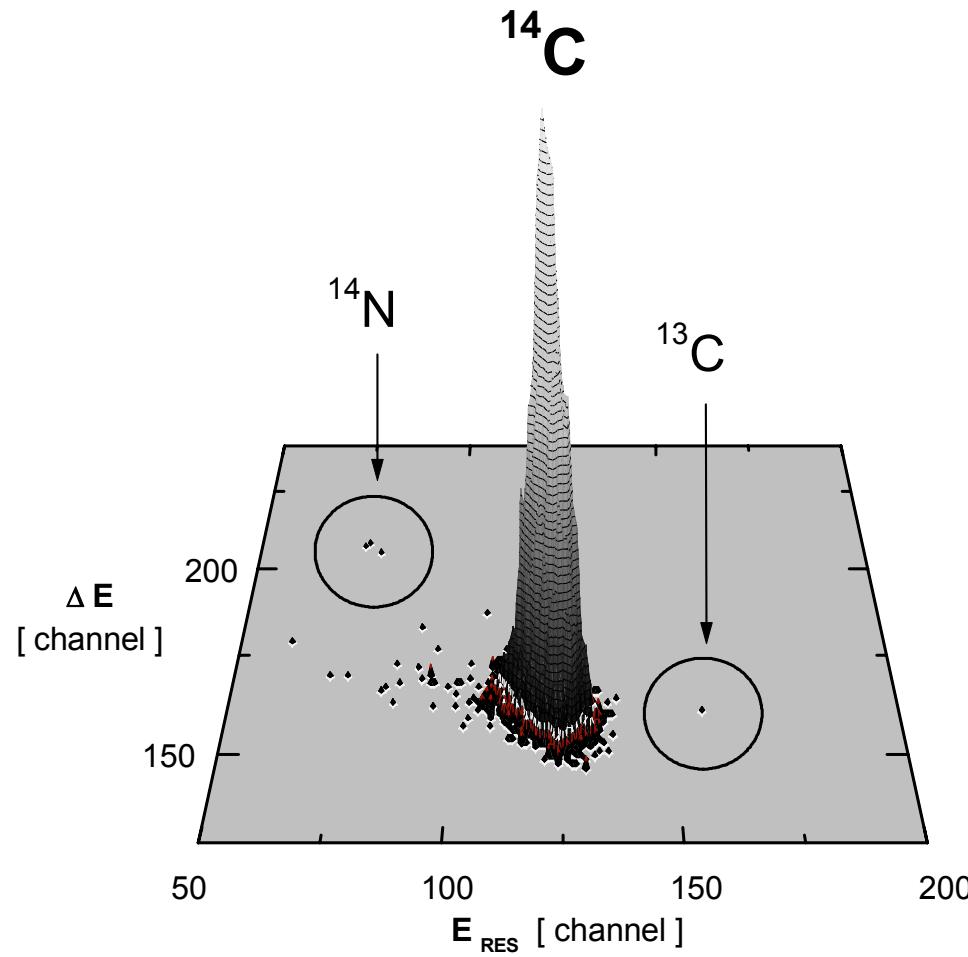
# Counting rare radionuclides before the invention of AMS



# Accelerator Mass Spectrometry



# Atom counting



# AMS detection of long-lived cosmogenic RN



**Table 1.1** Long-Lived Radioisotopes Routinely Measured by AMS

Radioisotope	$^{10}\text{Be}$	$^{14}\text{C}$	$^{26}\text{Al}$	$^{36}\text{Cl}$	$^{41}\text{Ca}$	$^{129}\text{I}$
Half-life (a)	1.51 M	5.73 k	720 k	301 k	103 k	15.7 M
Stable isotopes	$^9\text{Be}$	$^{12,13}\text{C}$	$^{27}\text{Al}$	$^{35,37}\text{Cl}$	$^{40,42,43,44}\text{Ca}$	$^{127}\text{I}$
Stable isobars	$^{10}\text{B}$	$^{14}\text{N}^{\text{a}}$	$^{26}\text{Mg}^{\text{a}}$	$^{36}\text{Ar}^{\text{a}}, ^{36}\text{S}$	$^{41}\text{K}$	$^{129}\text{Xe}^{\text{a}}$
Chemical form	$\text{BeO}$	$\text{C}$	$\text{Al}_2\text{O}_3$	$\text{AgCl}$	$\text{CaH}_2/\text{CaF}_2$	$\text{AgI}$
Sample size (mg)	0.5	0.02–1	2	1	10	2
Sensitivity (atom ratio)	$2 \times 10^{-15}$	$1 \times 10^{-15}$	$2 \times 10^{-15}$	$1 \times 10^{-15}$	$5 \times 10^{-15}$	$5 \times 10^{-14}$

<sup>a</sup> Do not form negative ions.

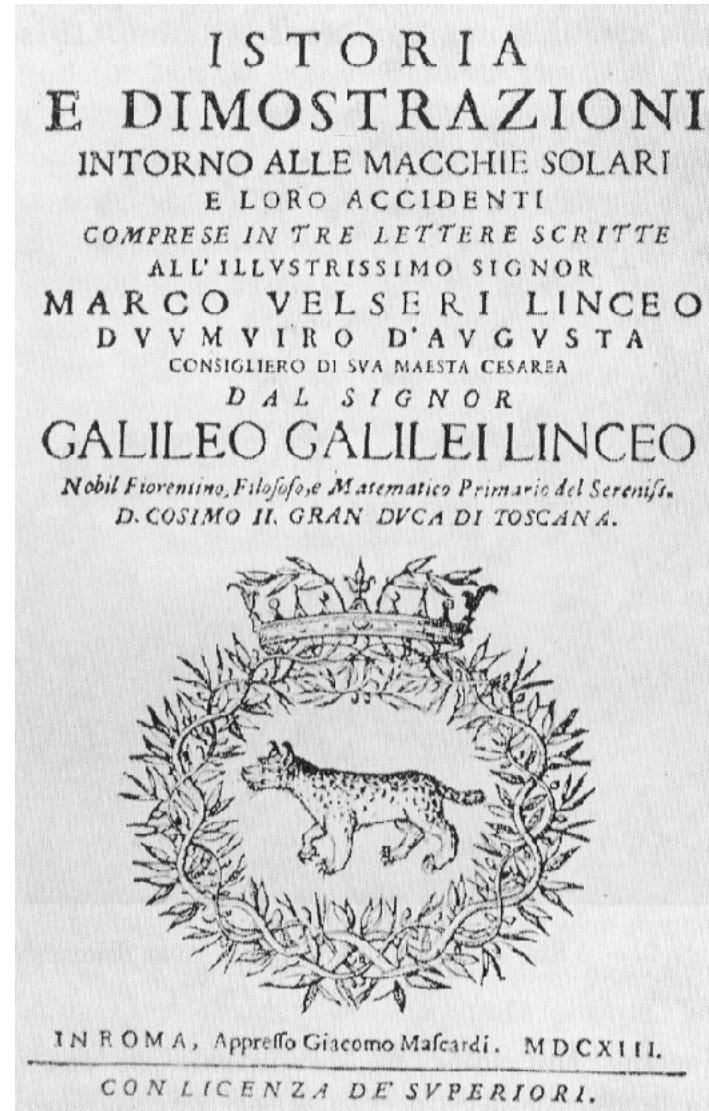
# Sun



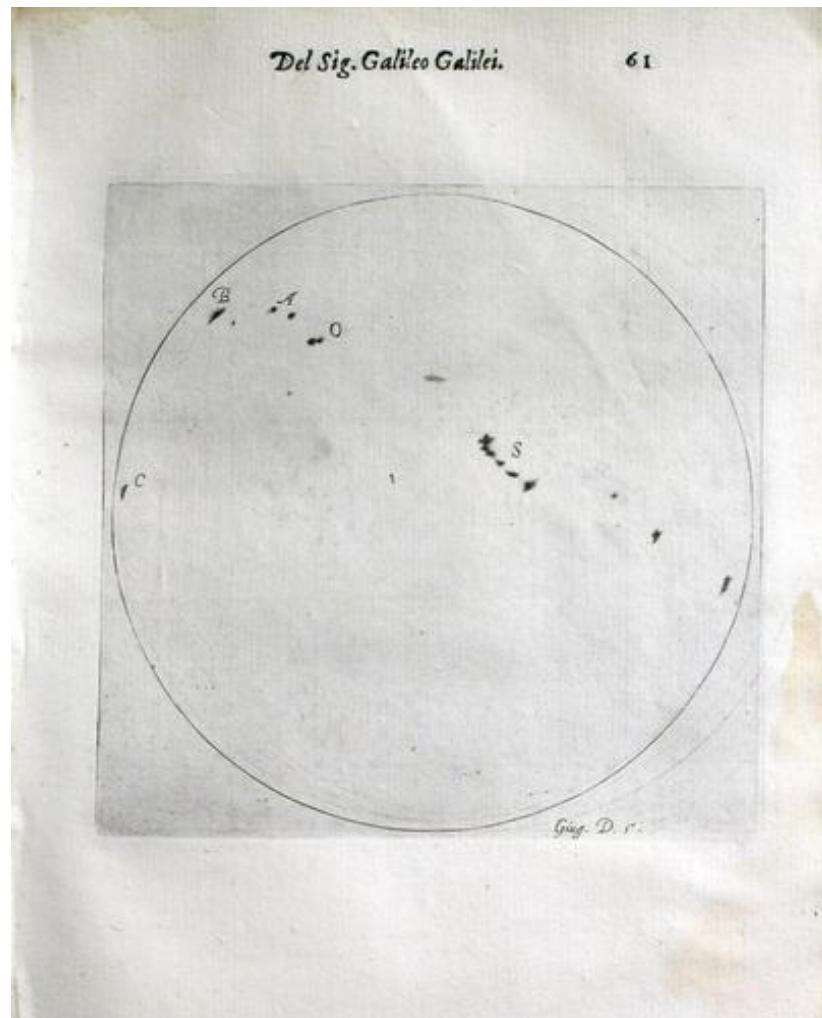
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Historical, instrumental and proxy record

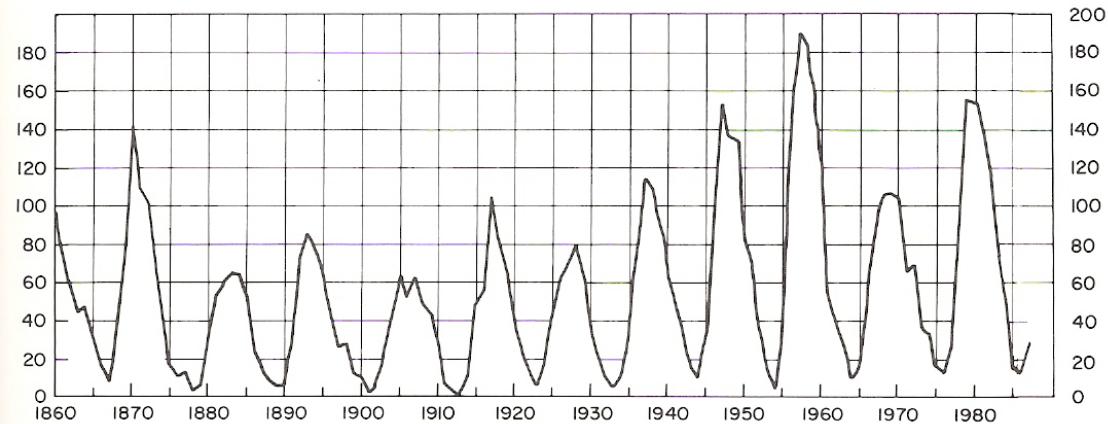
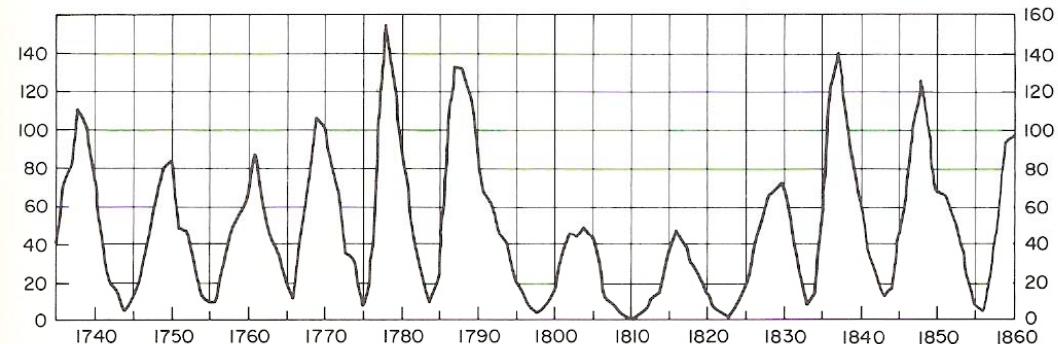
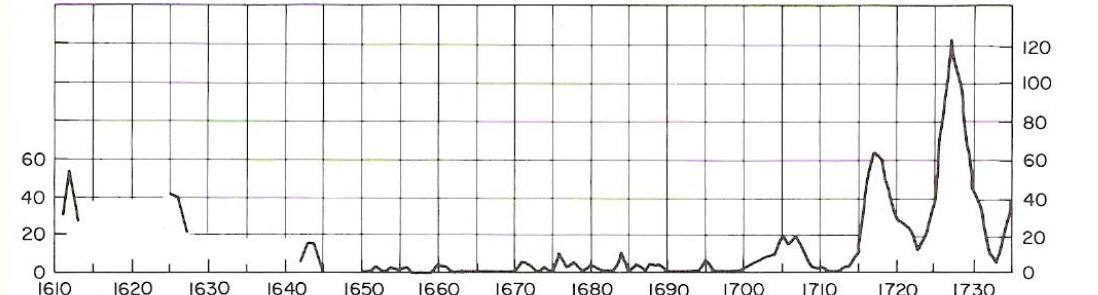
# HISTORICAL RECORD



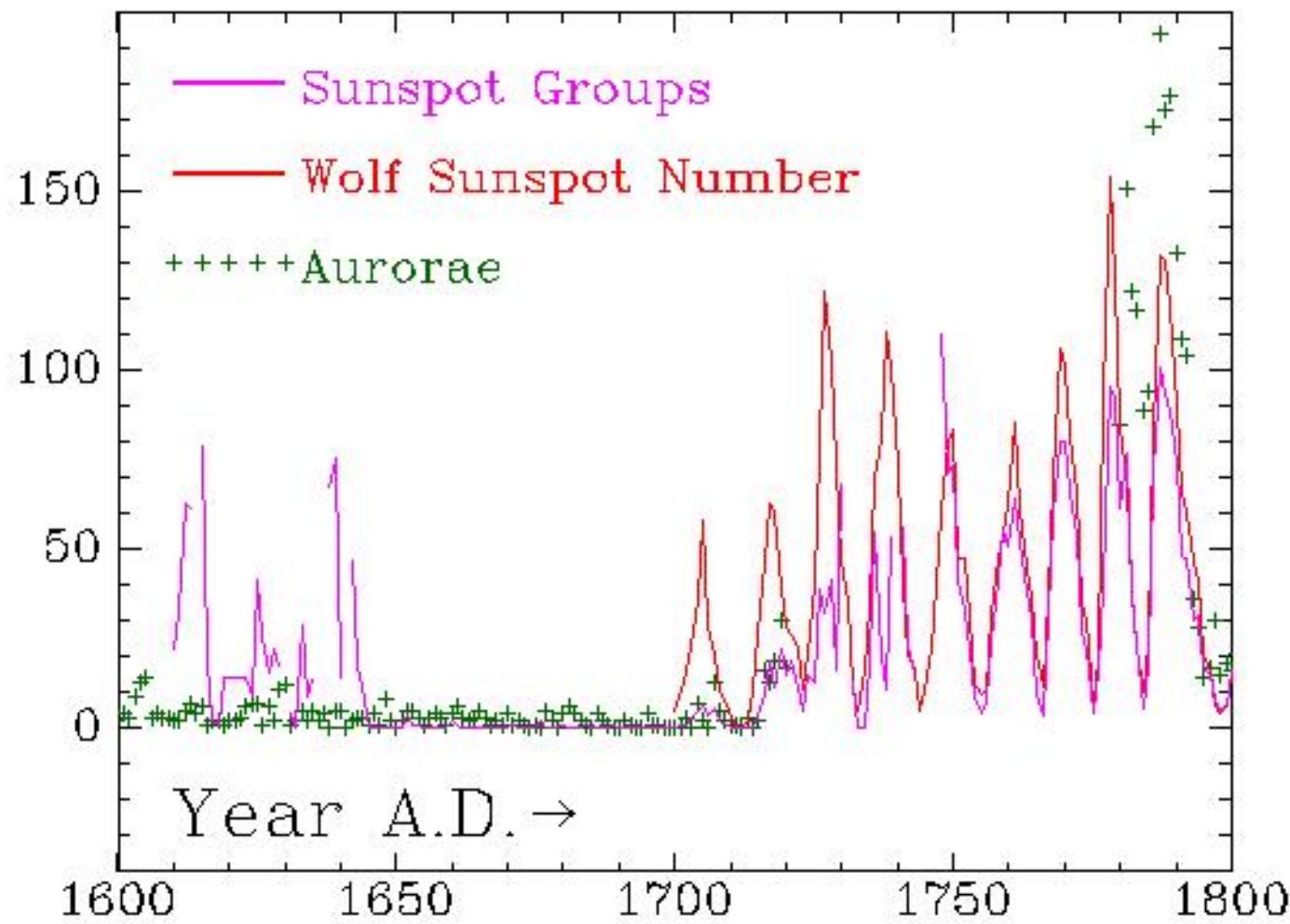
1613



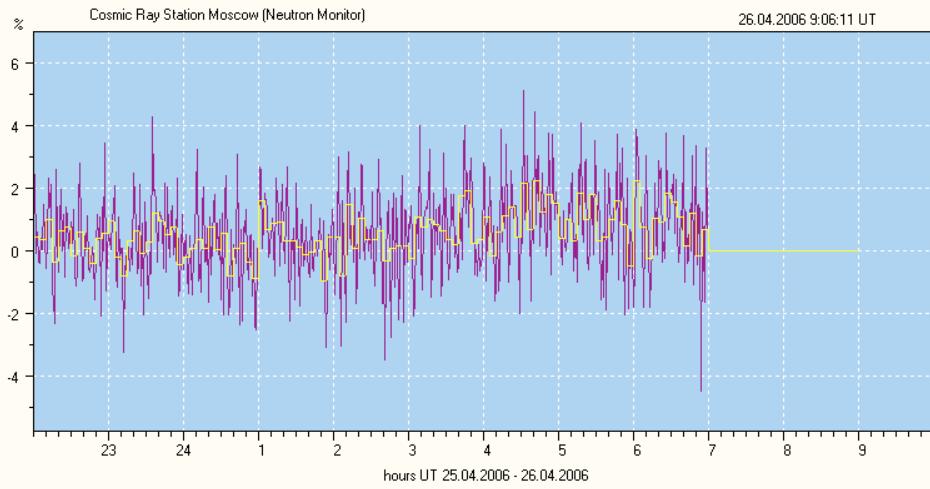
# Annual mean sunspot numbers 1610-1987



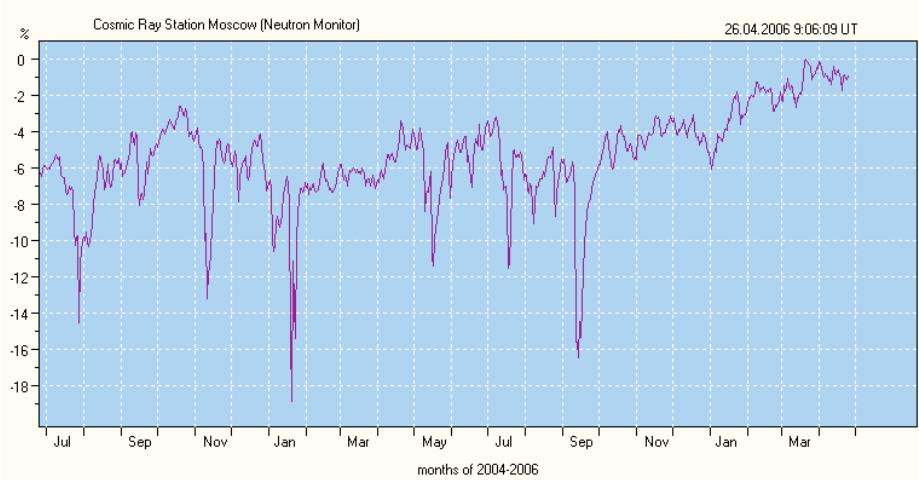
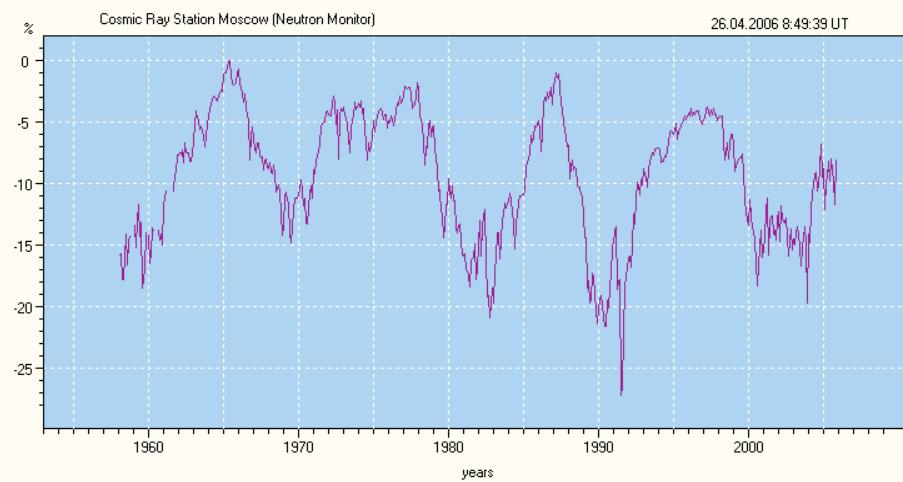
Cosmic Rays, M.W. Friedlander



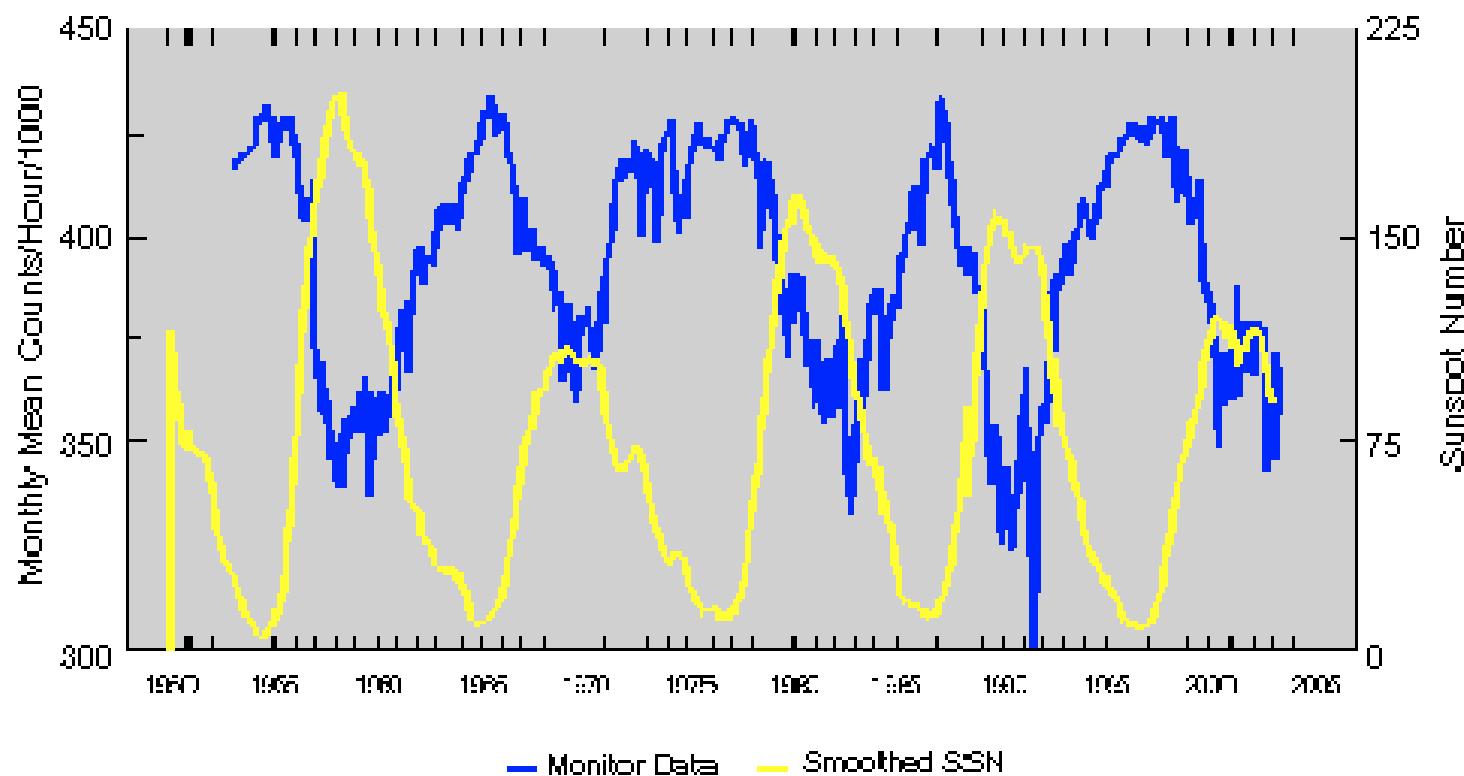
# INSTRUMENTAL RECORD



***Moscow Neutron Monitor***



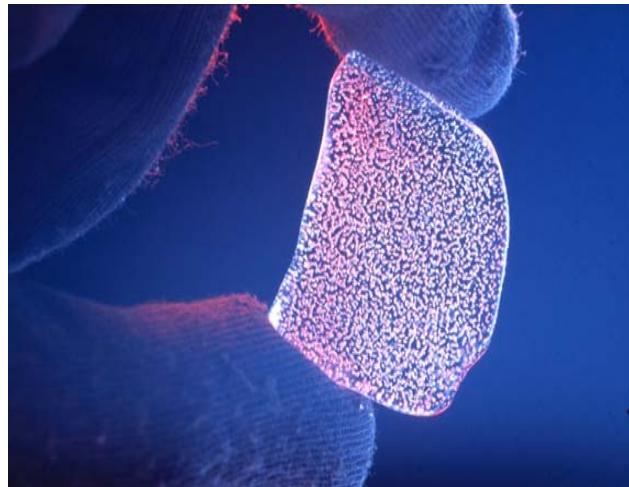
### Climax Corrected Neutron Monitor Values Smoothed Sunspot Numbers 1950-2002



# PROXY RECORD

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Tree Rings, Ice Cores, Sediments, Varves and Corals



# $^{10}\text{Be}$ and $^{14}\text{C}$ production

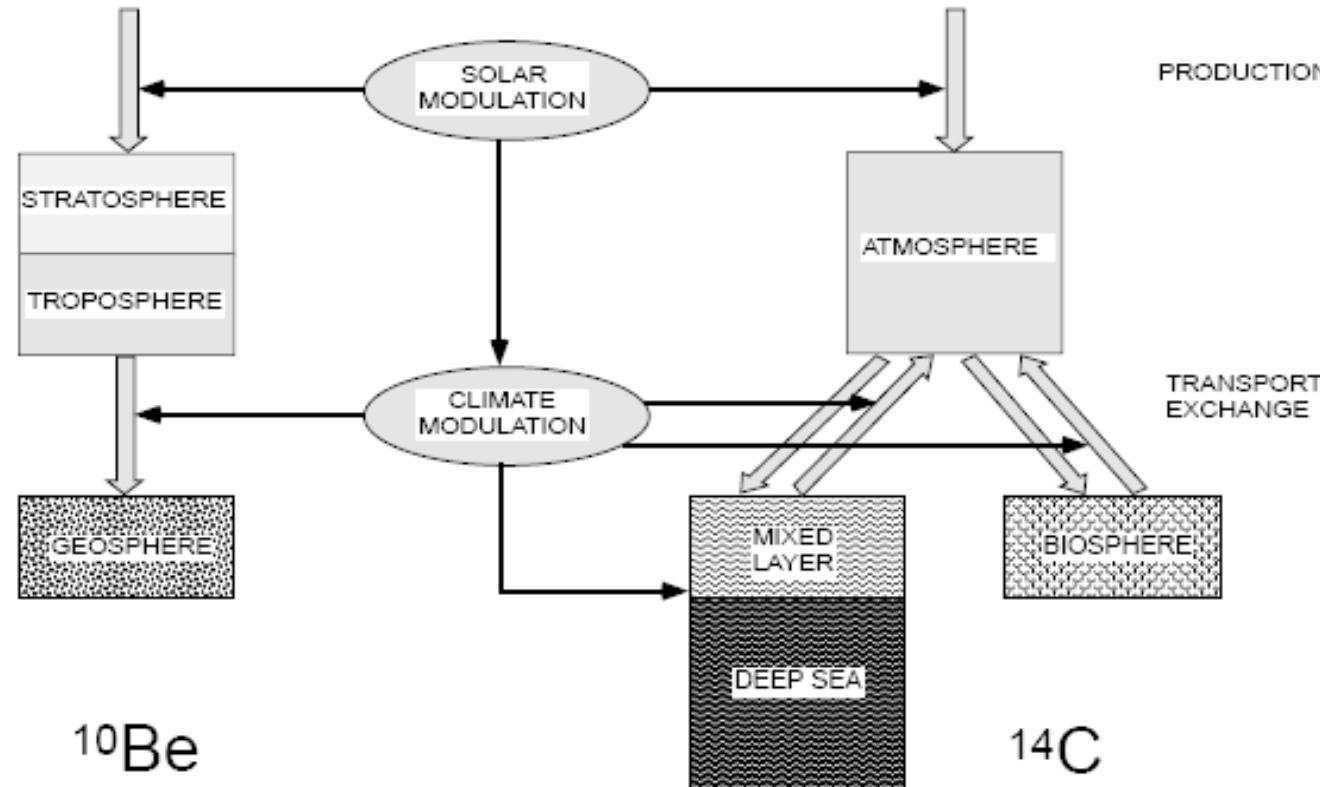


Figure 4. Description of the  $^{10}\text{Be}$  and  $^{14}\text{C}$  systems by simple box models. Solar modulation affects the production rate of both nuclides in a very similar way. Climate modulation (transport, atmospheric mixing, and deposition), however, leads to completely different effects in the two systems. After its removal from the atmosphere  $^{10}\text{Be}$  becomes stored in the geosphere. In contrast  $^{14}\text{C}$  continuously exchanges between the different reservoirs resulting in a much more complex behaviour.

Beer, Space Science Reviews, 2000

# Long-lived cosmogenic radionuclides



Nuclide	Target	Half-life [y]	Production Rate [atoms cm <sup>-2</sup> s <sup>-1</sup> ]	Inventory [tons]
<sup>14</sup> C	N	5730	2	62
<sup>10</sup> Be	N, O	$1.51 \times 10^6$	0.018	105
<sup>36</sup> Cl	Ar	$3.08 \times 10^5$	0.0019	8

Beer, Space Science Reviews, 2000

# $^{10}\text{Be}$ production

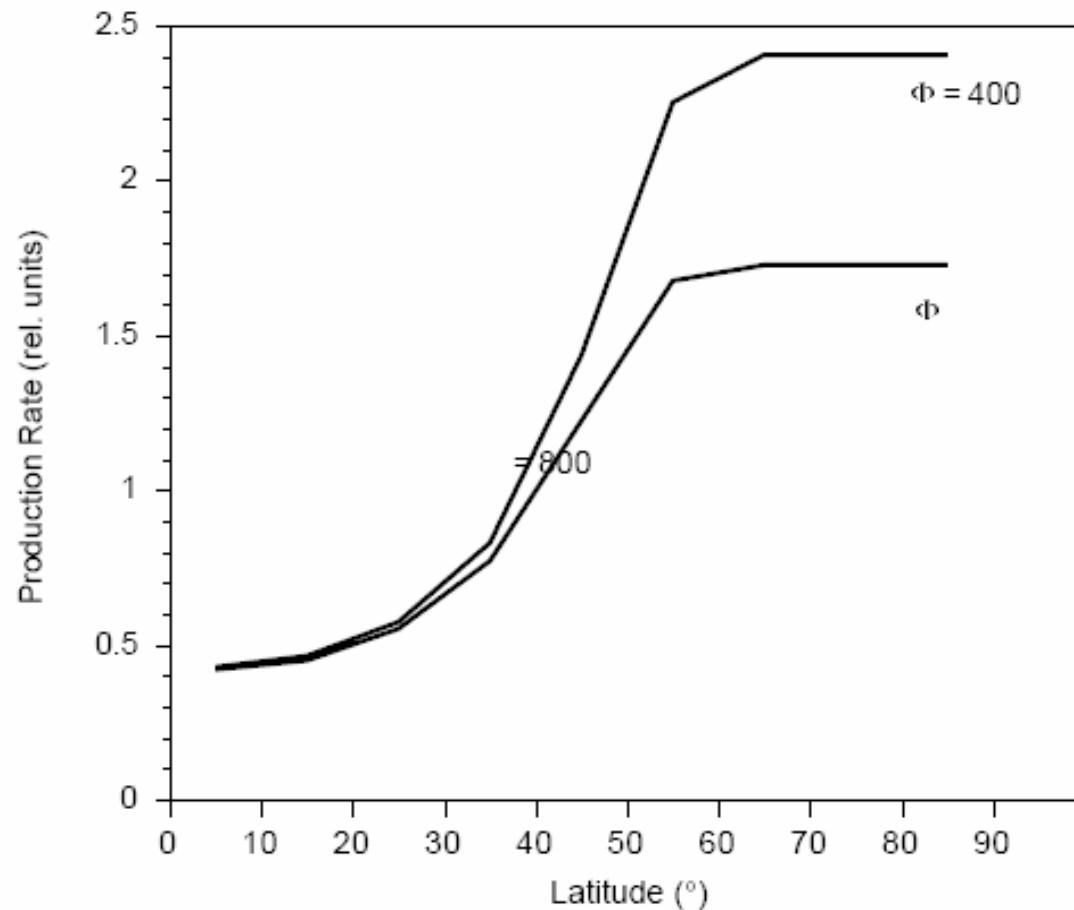


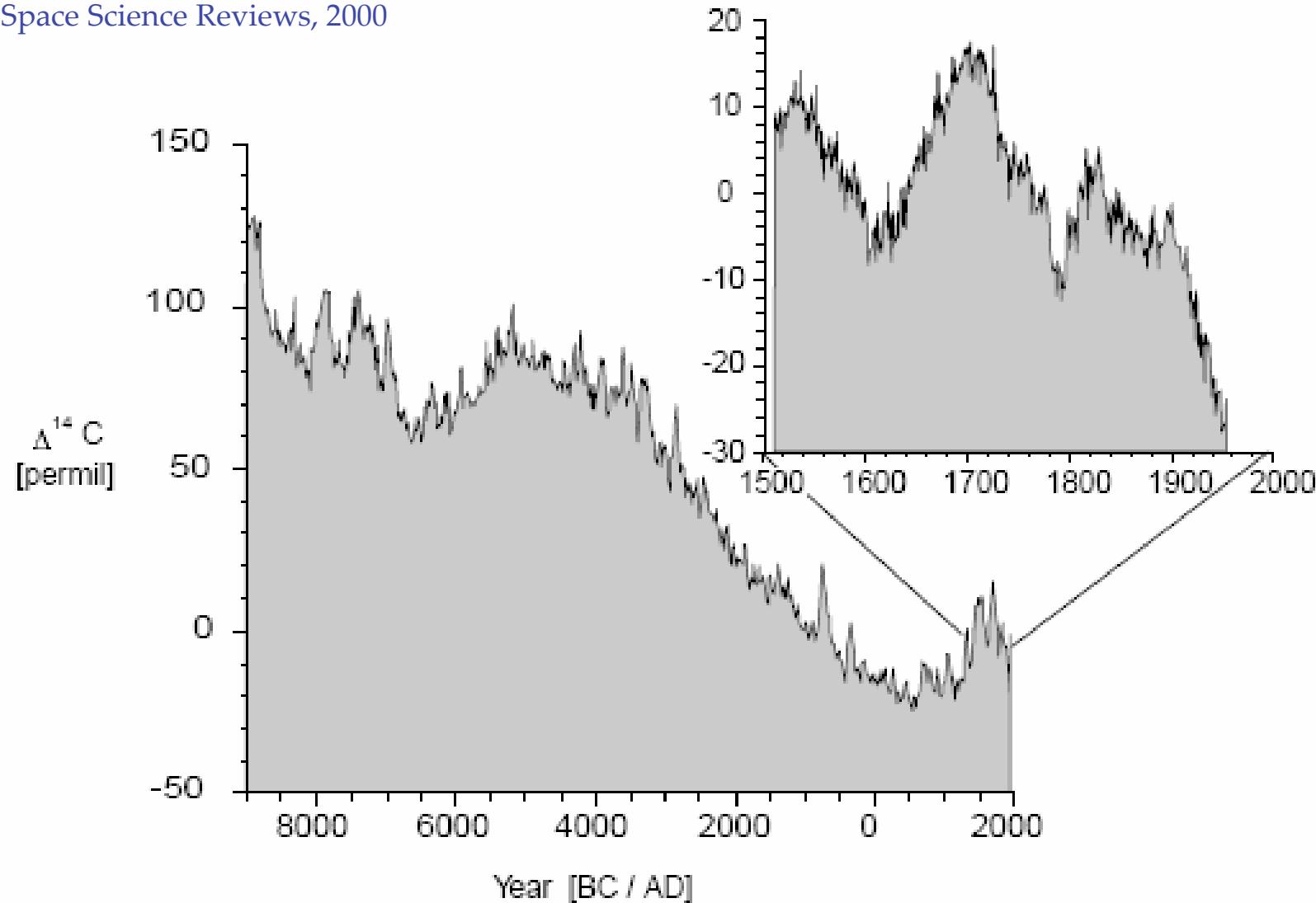
Figure 3. Relative production rate of  $^{10}\text{Be}$  as a function of geomagnetic latitude for the present geomagnetic field intensity ( $M = 1$ ) and two different solar activity levels. Note that the solar modulation is effective mainly at high latitudes.

Beer, Space Science Reviews, 2000

# $^{14}\text{C}$ in tree rings



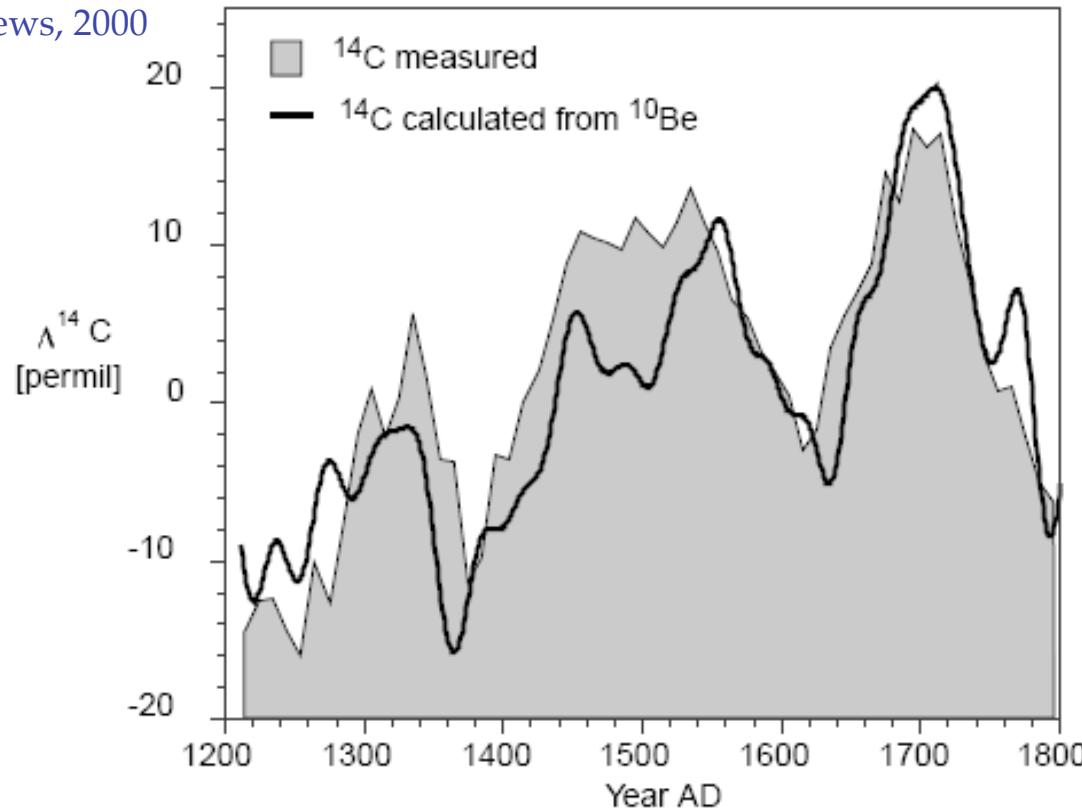
Beer, Space Science Reviews, 2000



# $^{10}\text{Be}$ and $^{14}\text{C}$ record

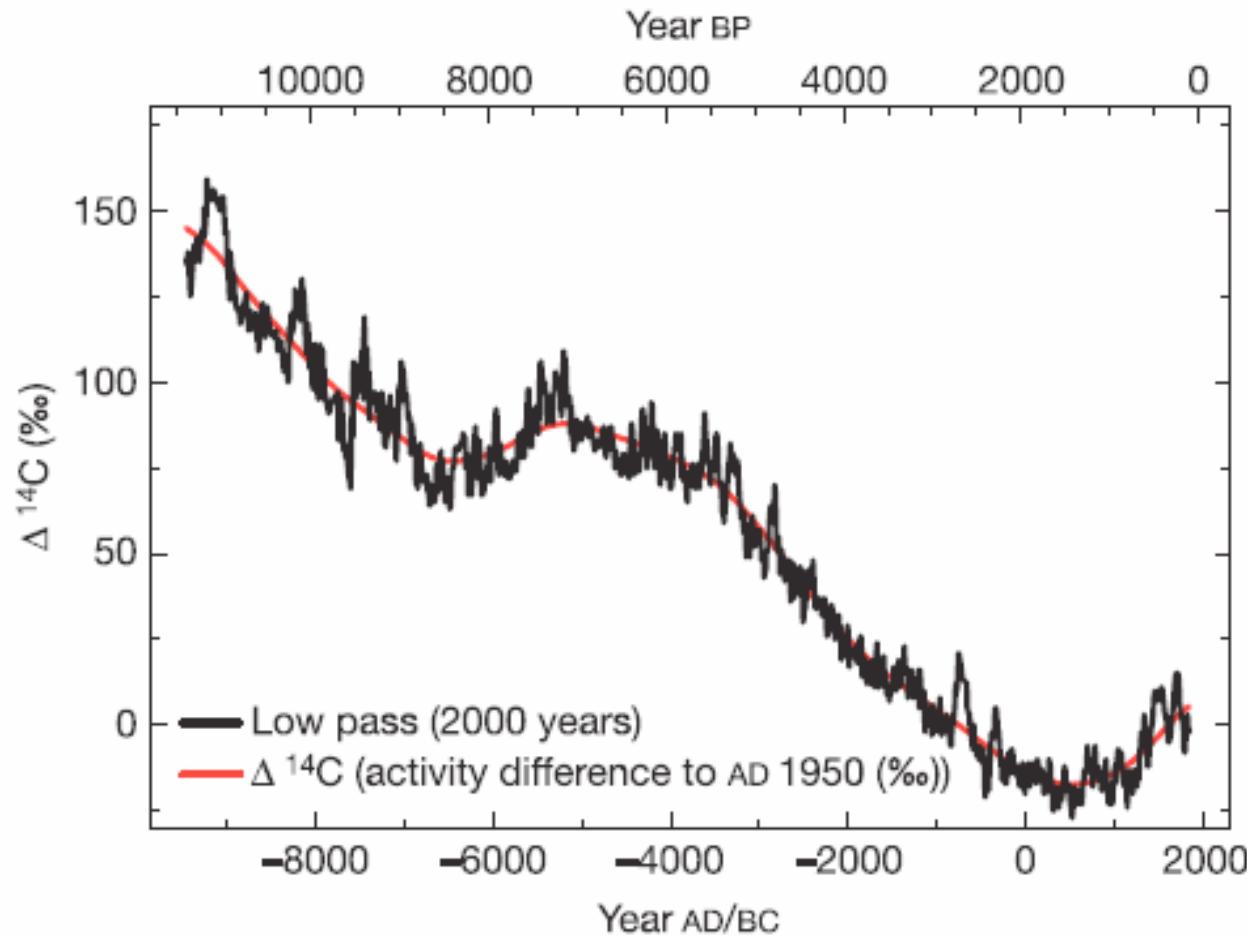


Beer, Space Science Reviews, 2000



*Figure 7.* Comparison of the  $\Delta^{14}\text{C}$  record measured on tree rings with the  $\Delta^{14}\text{C}$  record calculated based on a combined  $^{10}\text{Be}$  record of two ice cores from Milcent (Greenland) and the South Pole. The calculation was performed using the carbon cycle model of Fig. 4 assuming that the combined  $^{10}\text{Be}$  record reflects directly the atmospheric production rate. The three periods of low solar activity (Maunder, Spoerer and Wolf) are clearly visible.

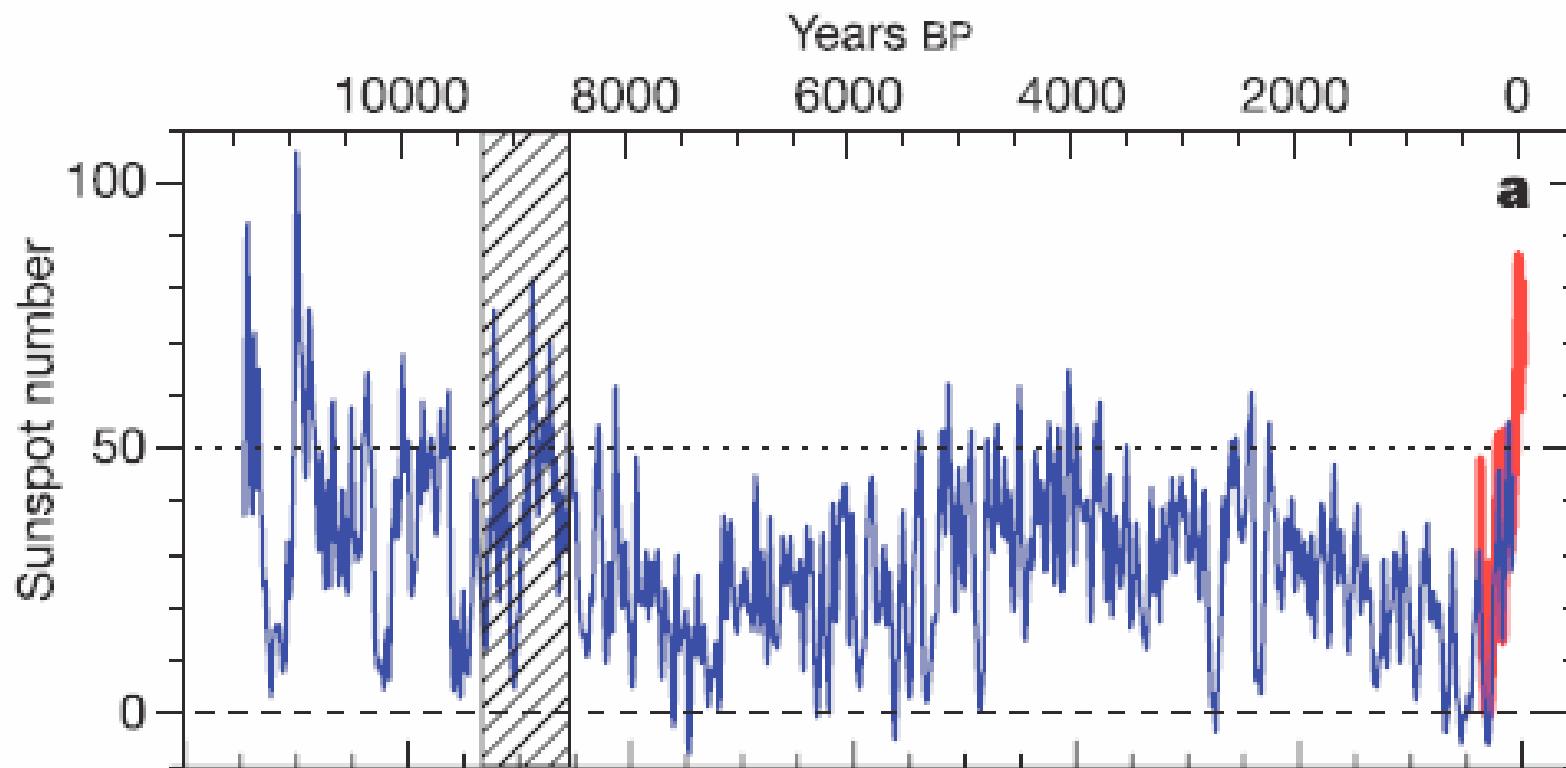
# $^{14}\text{C}$ in tree rings



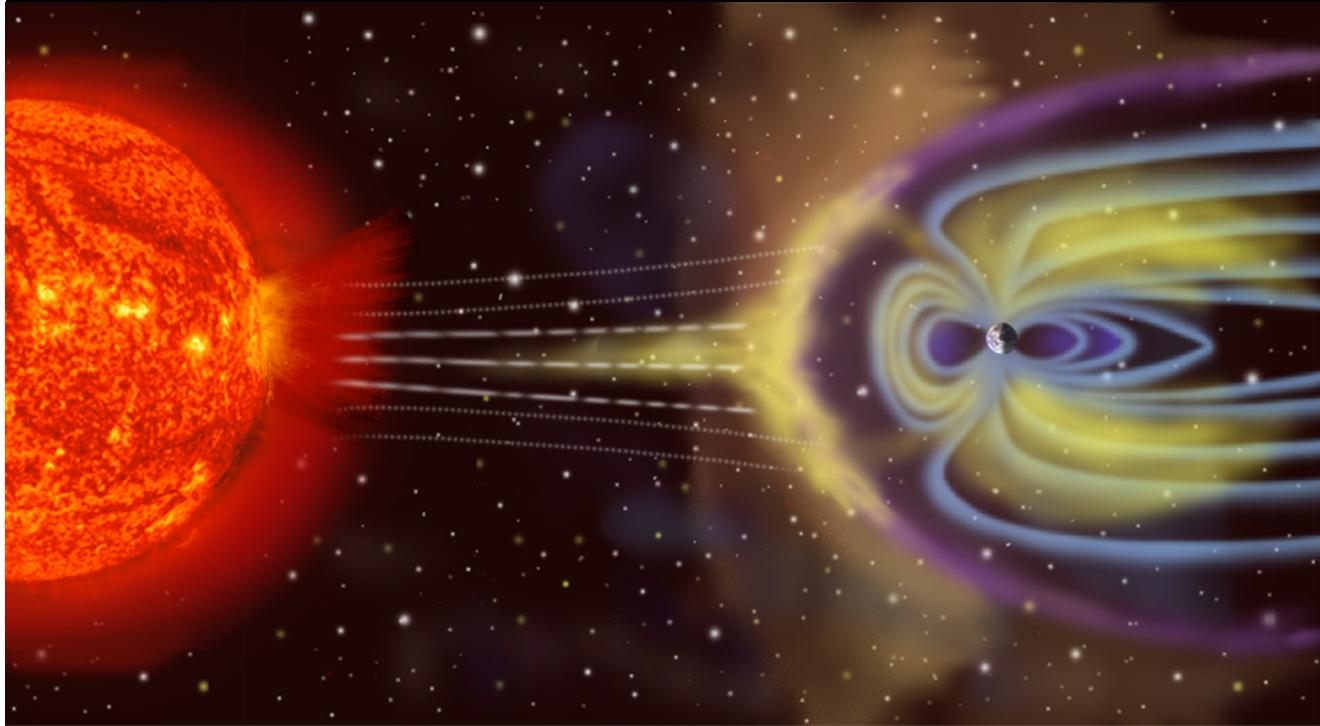
# Unusual activity of the Sun during recent decades compared to the previous 11,000 years

S. K. Solanki<sup>1</sup>, I. G. Usoskin<sup>2</sup>, B. Kromer<sup>3</sup>, M. Schüssler<sup>1</sup> & J. Beer<sup>4</sup>

Nature 2004



# $^{10}\text{Be}$ from the Sun



$^{10}\text{Be}$  in excess of that expected from in situ cosmic ray spallation reactions is present in lunar surface soil 78481; This excess  $^{10}\text{Be}$  and its association with surficial layers corresponds to  $1.9 \times 10(8)$  atoms per square centimeter, requiring a  $^{10}\text{Be}$  implantation rate of  $2.9 \times 10(-6)$  atoms per square centimeter per second on the surface of the Moon. The most likely site for the production of this excess  $^{10}\text{Be}$  is the Sun's atmosphere. The  $^{10}\text{Be}$  is entrained into the solar wind and transported to the lunar surface



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**Thanks!**