



The Abdus Salam
International Centre for Theoretical Physics



310/1749-10

**ICTP-COST-CAWSES-INAF-INFN
International Advanced School
on
Space Weather
2-19 May 2006**

***The Formation of the Earth's
Magnetosphere:
Basic physical principles and steady-
state configuration***

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These lecture notes are intended only for distribution to participants

The Formation of the Earth's Magnetosphere: Basic physical principles and steady-state configuration



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May 2006



*The support provided by COST, COST Action 724
and the Abdus Salam ICTP is gratefully acknowledged*

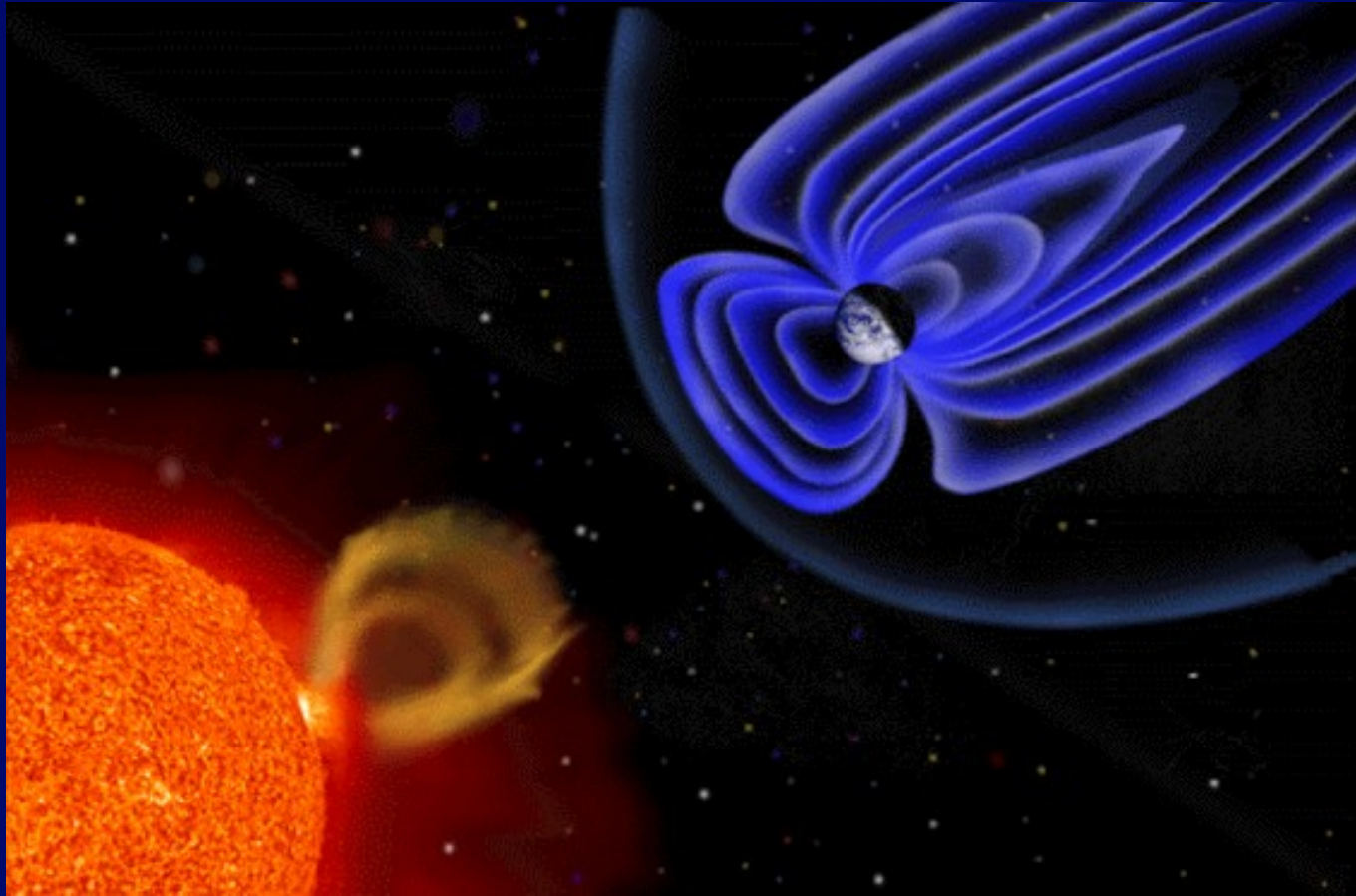


Figure by Steele Hill (from D. Stern "The exploration of the Earth's magnetosphere")

The formation of a planetary magnetosphere requires:

- a quasi-permanent magnetic field of the planet
- an ionized plasma stream which interacts with the planetary magnetic field and thereby shapes its "magnetosphere"

What is a plasma ?

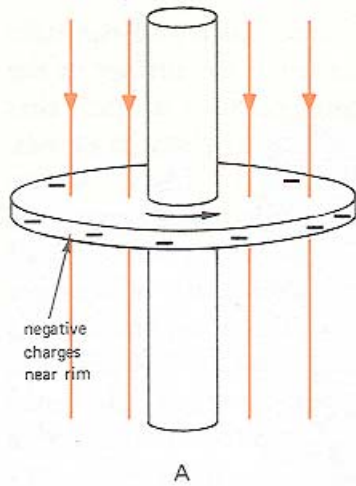
Matter which is ionized to such an extent that the electromagnetically dominated collective behavior of the electrons and ions governs the dynamics of the medium. The solar wind, the ionized upper atmosphere and the fluid outer core of the Earth are examples of matter in the plasma state.

The troposphere contains ions (produced by cosmic rays, UV radiation, candle light, ...) but the charged particle density is too small compared to the neutral air to enable collective behavior. The relatively low charged particle density allows only single particle behavior dominated by mechanical forces (gravity, collisions).

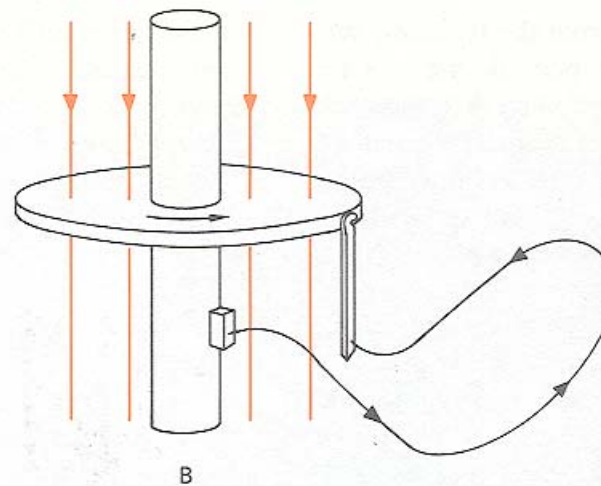
- Planet Earth has a quasi-static magnetic field of internal origin (30 - 60 μT at the Earth's surface)
- Mercury has a (weak) magnetic field
- Mars and Venus have at present time no measurable magnetic field (Mars has some remnants), consequently they do not possess a magnetosphere
- The giant gas planets and the outer planets are known to have a magnetic field

The magnetic field measured at the surface of the Earth is too strong to be generated solely by (crustal) permanent magnetization.
It must have a different source

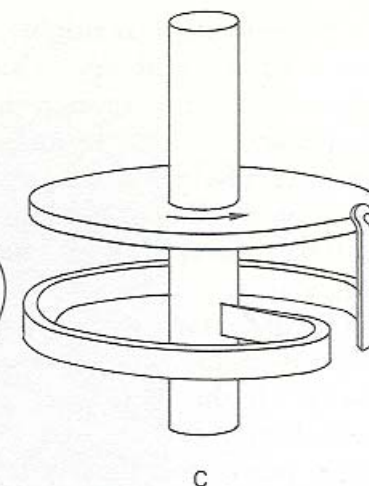
The geodynamo



4.7 (a) Disc rotating in magnetic field, no current flows.



(b) External circuit enables current to flow.



(c) Current provides a field and gives a self-exciting dynamo.

electrically conducting disc:
 $\mathbf{v} \times \mathbf{B}$ force drives positively charged particles toward the center

Figure from E. Bullard, "The Earth's magnetic field and its origin"

Cowling theorem (1933):

An axisymmetric magnetic field generated by currents at finite distance cannot be maintained by a plasma velocity field of finite magnitude

The way out: Three-dimensional "mean field" electrodynamics

$$\mathbf{v} = \bar{\mathbf{v}} + \mathbf{v}'$$

$$\mathbf{B} = \bar{\mathbf{B}} + \mathbf{B}'$$

$$\mathbf{j} = \bar{\mathbf{j}} + \mathbf{j}' = \sigma \{ \bar{\mathbf{E}} + \mathbf{E}' + (\bar{\mathbf{v}} + \mathbf{v}') \times (\bar{\mathbf{B}} + \mathbf{B}') \}$$

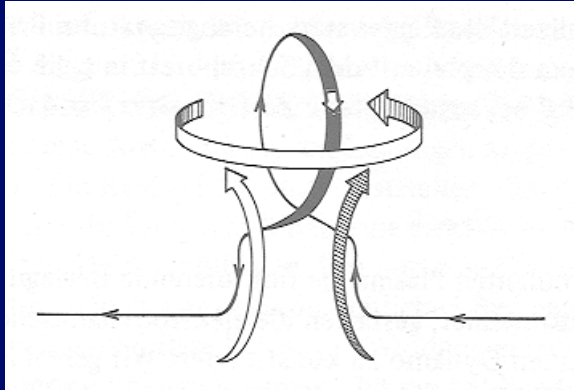
$$\bar{\mathbf{j}} = \sigma \{ \bar{\mathbf{E}} + \bar{\mathbf{v}} \times \bar{\mathbf{B}} + \overline{(\mathbf{v}' \times \mathbf{B}')} \}$$

\mathbf{v}' and \mathbf{B}' assumed to be correlated (and a few further assumptions) lead to

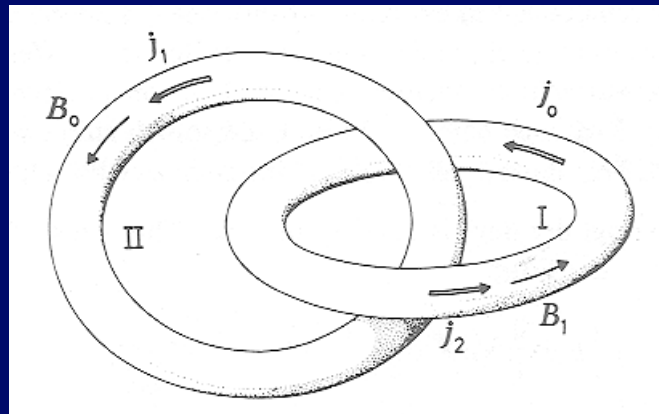
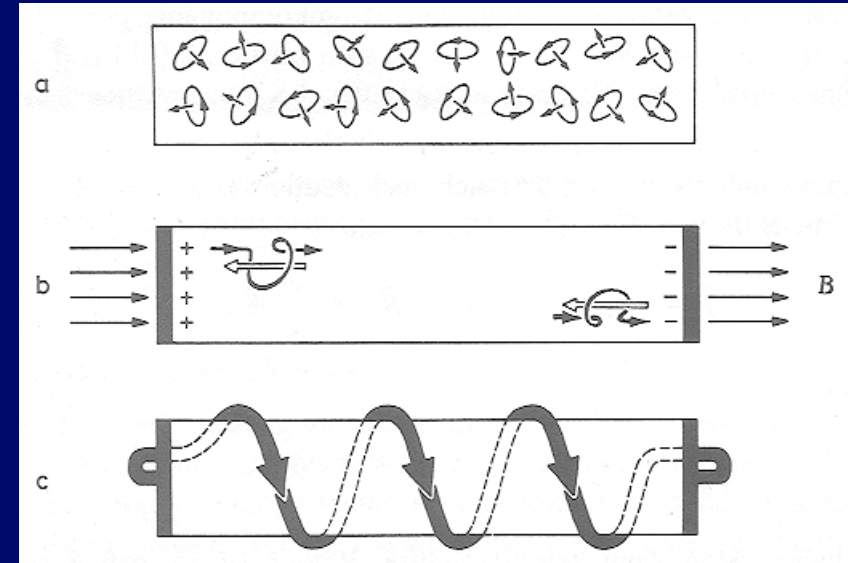
$$\overline{(\mathbf{v}' \times \mathbf{B}')} = \alpha \bar{\mathbf{B}} - \beta \nabla \times \bar{\mathbf{B}}$$

If turbulence has a preference rotation then

$$\alpha \neq 0 \Rightarrow \alpha\text{-effect}$$



perfectly conducting plasma:
magnetic field is "frozen-in"



Figures from Kippenhahn/Moellenhoff,
"Elementary plasma physics" (1975)

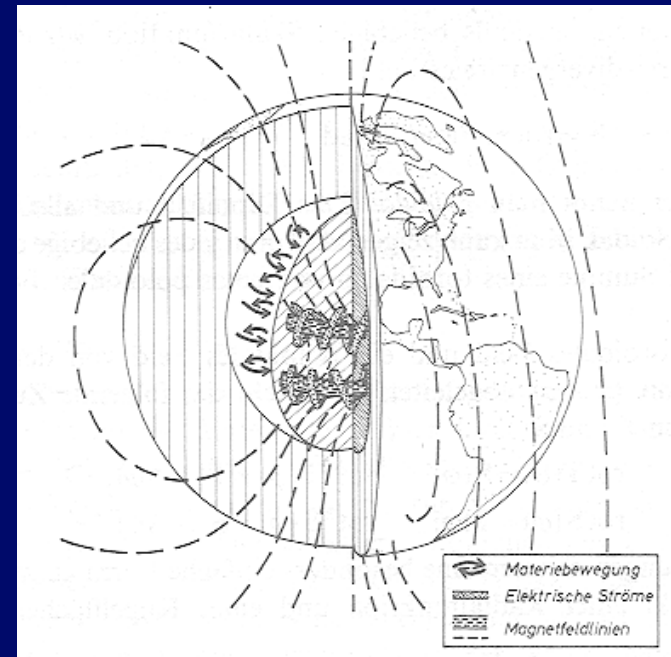


Figure from Kertz after Elsasser (1958)

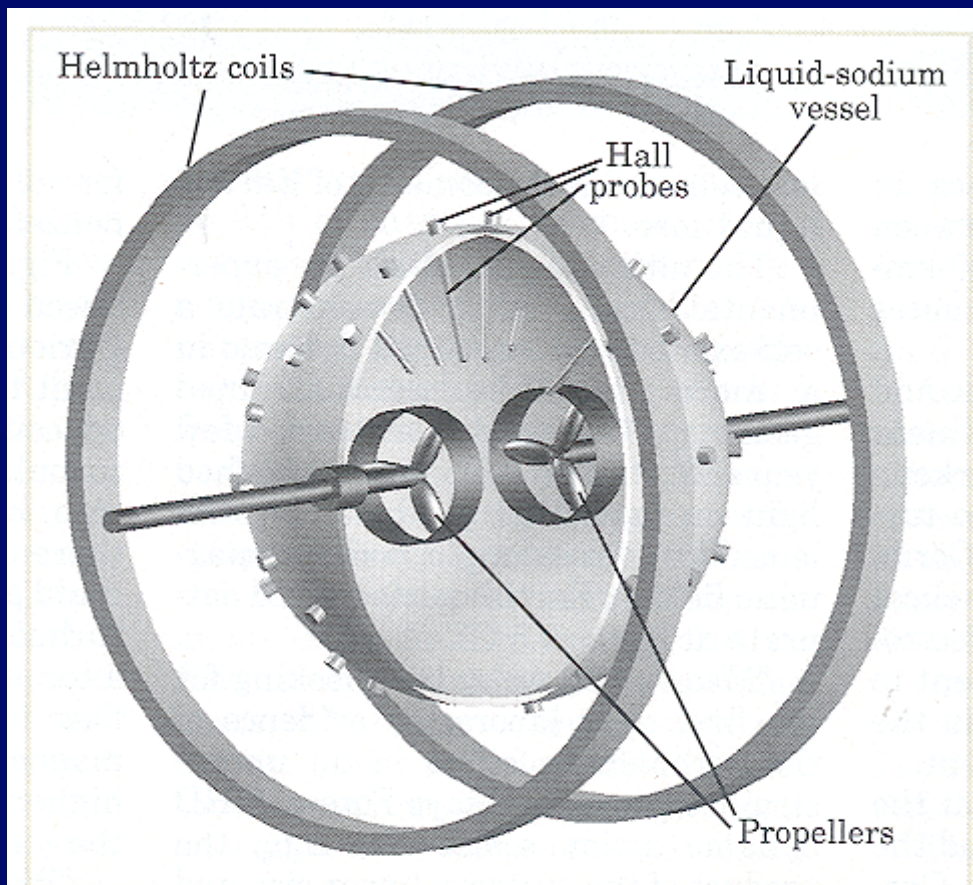


Figure 1. In the Madison Dynamo Experiment, liquid sodium filling a 1-meter-diameter spherical vessel is agitated by two counterrotating propellers. Helmholtz coils, coaxial with the propeller shafts, impose an almost uniform axial magnetic field on the vessel. Arrays of Hall probes measure the resultant magnetic field inside the liquid and on the vessel's surface. (Adapted from ref. 2.)

Laboratory dynamo experiment
Figure from "Physics Today", 2006

The solar wind

Flow of mainly protons and electrons

+ a few percent fully ionized Helium

+ a few heavier ions

(almost always supersonic and super-Alfvénic)

dragging along the interplanetary magnetic field

Solar wind speed derived from isothermal coronal expansion models

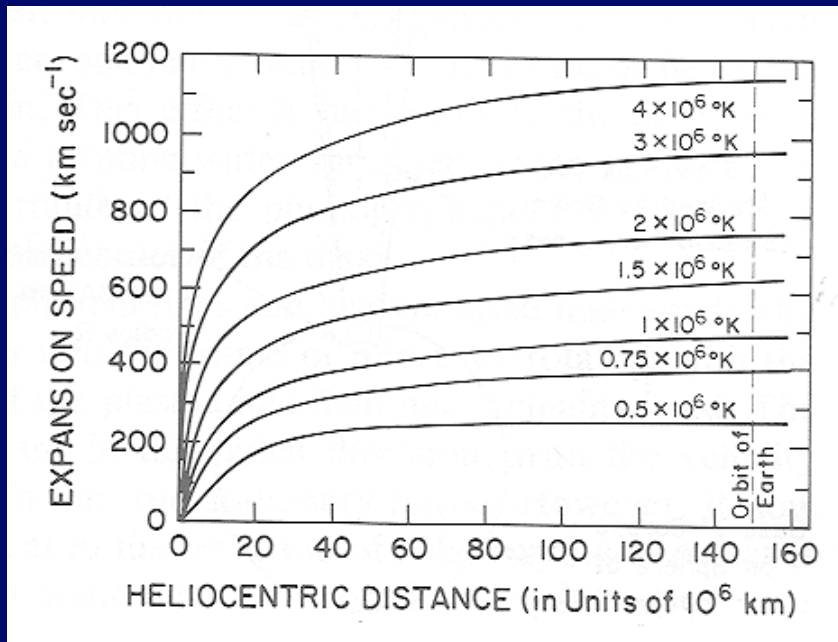


Figure after Parker (1958)

Solar wind speed inferred from radio wave scintillation observations

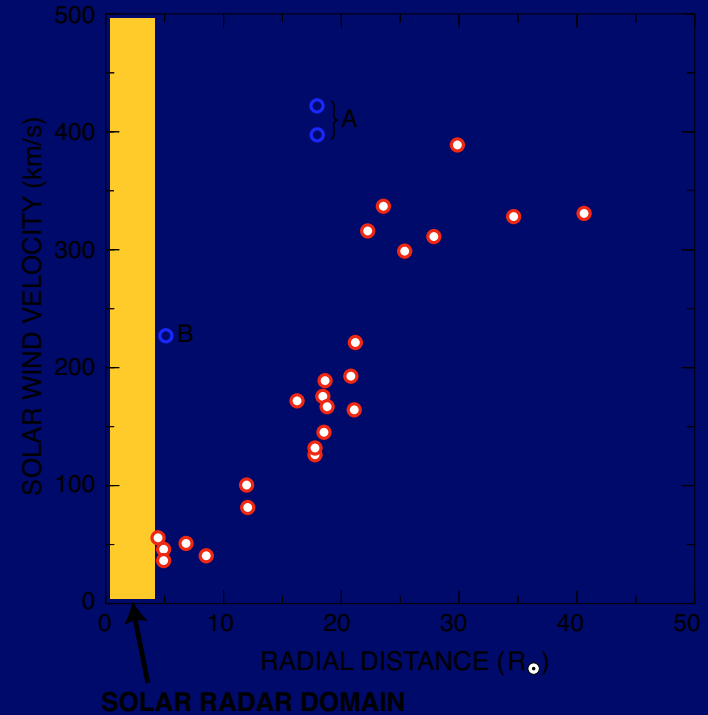
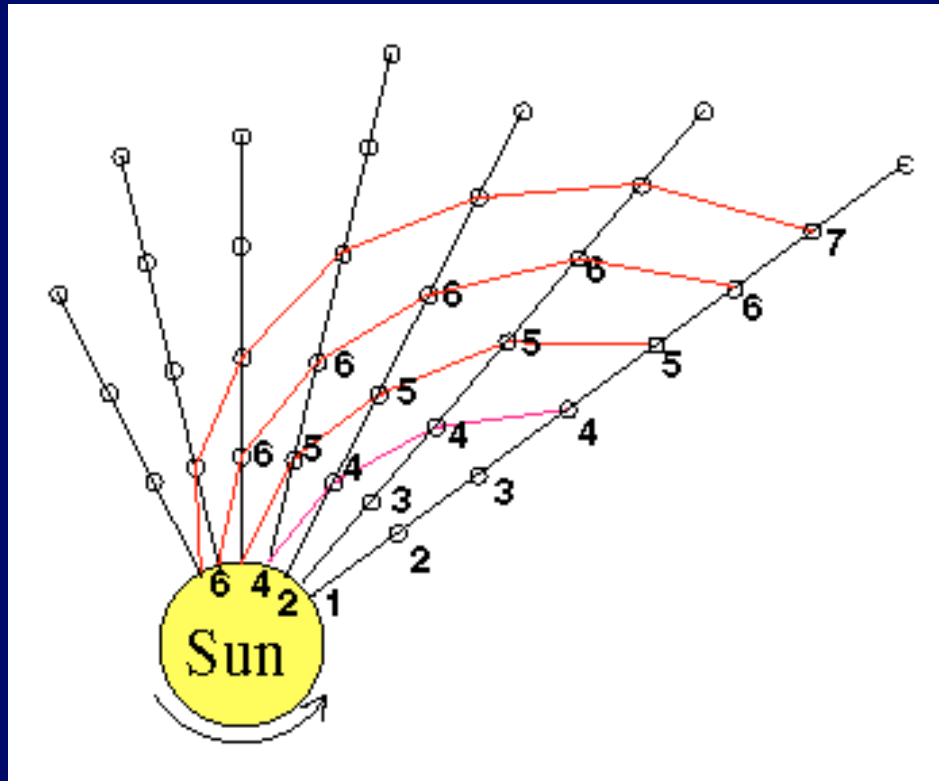


FIGURE 15: The solar wind velocity as a function of radial distance from the Sun obtained from interplanetary scintillation (IPS) measurements made at the Kashima 34-m radio telescope in Japan in 1989. The points indicated by A and B are thought to be related to coronal mass ejection (CME) events. A powerful solar radar would probably allow measurements of the early stages of solar wind acceleration and CME development at low solar altitudes. After Tokomaru *et al.* [108]; see also [109] and [120].

From "Radio studies of solar-terrestrial relationships"
The LOIS Science Team (2002)



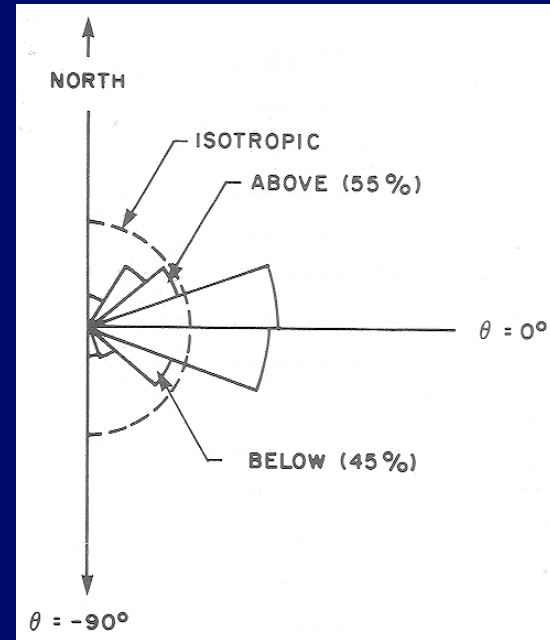
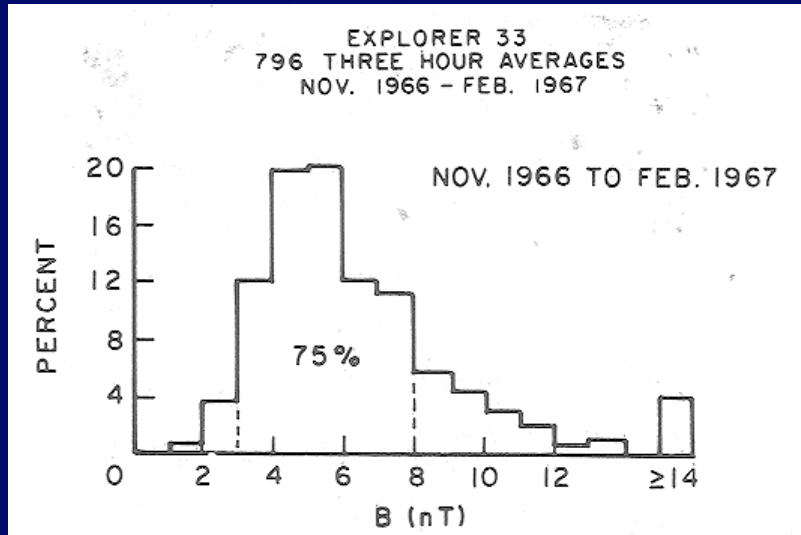
*Figure from D. Stern
"The exploration of the Earth's
Magnetosphere"*

Solar wind plasma expands radially outward

solar magnetic field is carried along ("frozen-in")

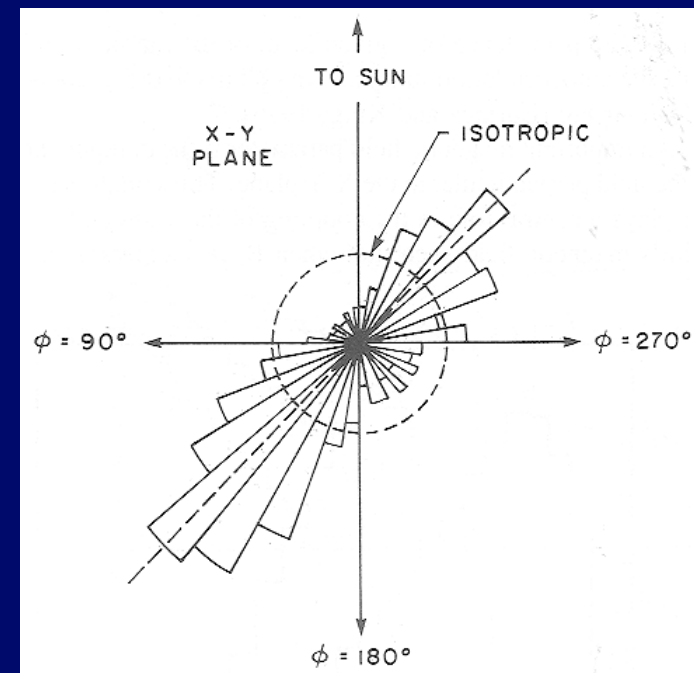
⇒ magnetic field lines form Archimedian spiral (Parker spiral)

at 1 AU (Earth orbit) $\approx 45^\circ$



Amplitude and orientation
of the interplanetary
magnetic field near 1 AU

Figures after Hirshberg, 1969



The current sheet which generates the sector structure observed near Earth

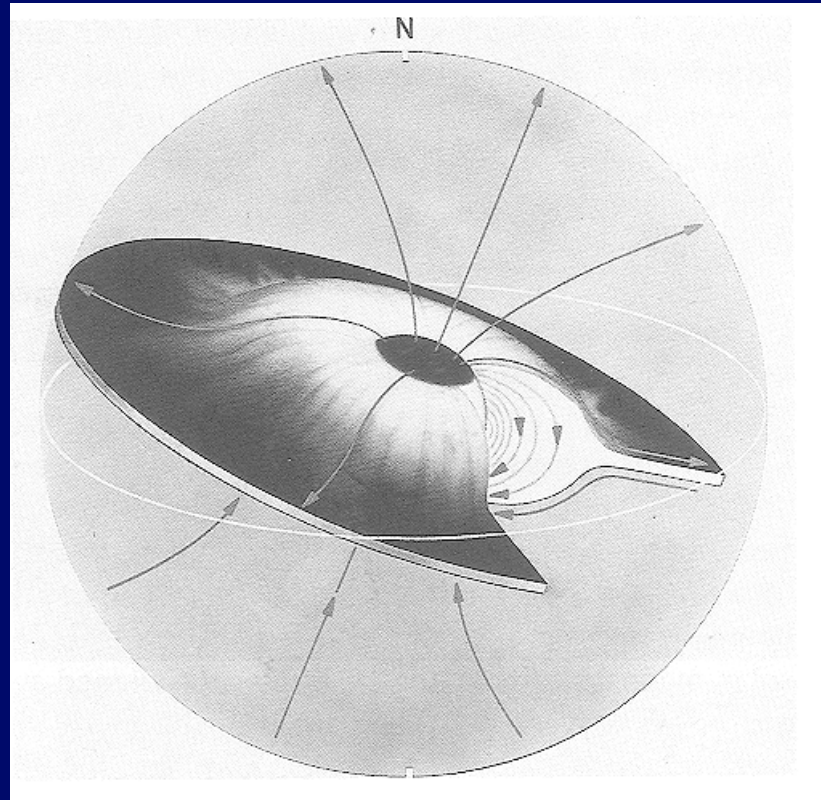


Figure from Hundhausen (1977)

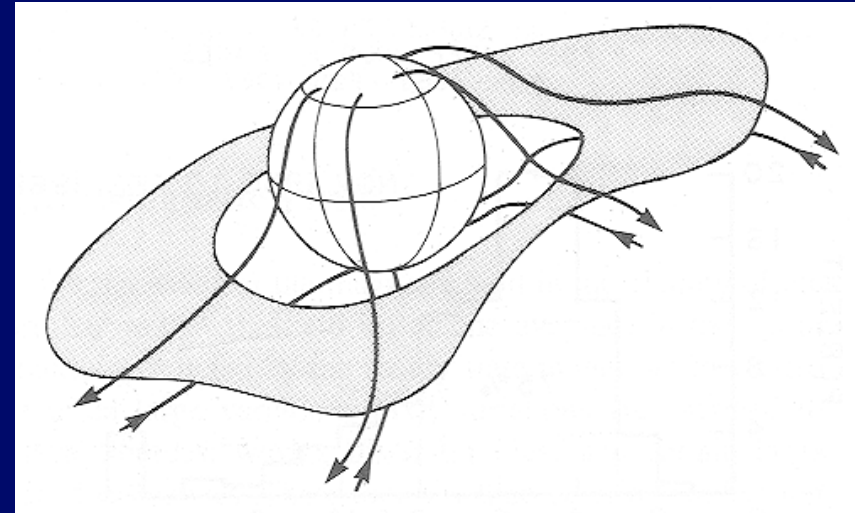


Figure from Smith (1979)

The solar wind

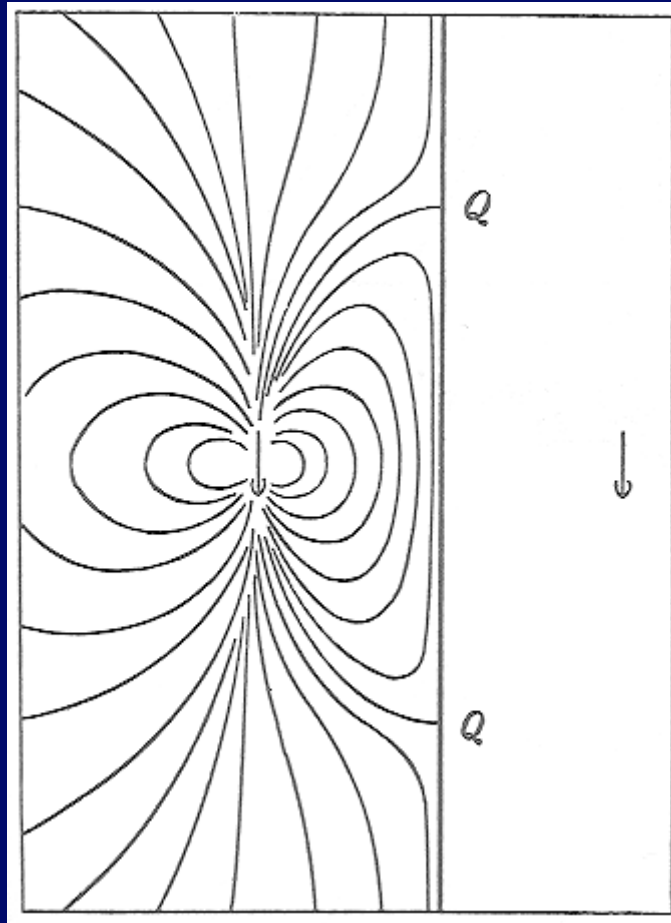
Typical numbers at 1 AU (slow solar wind)

plasma density	5-10 ions / cm ³
bulk speed	350-450 km/s
Mach number	6
p ⁺ and e ⁻ temperatures	1-2×10 ⁵ K
magnetic induction	5-10 nT

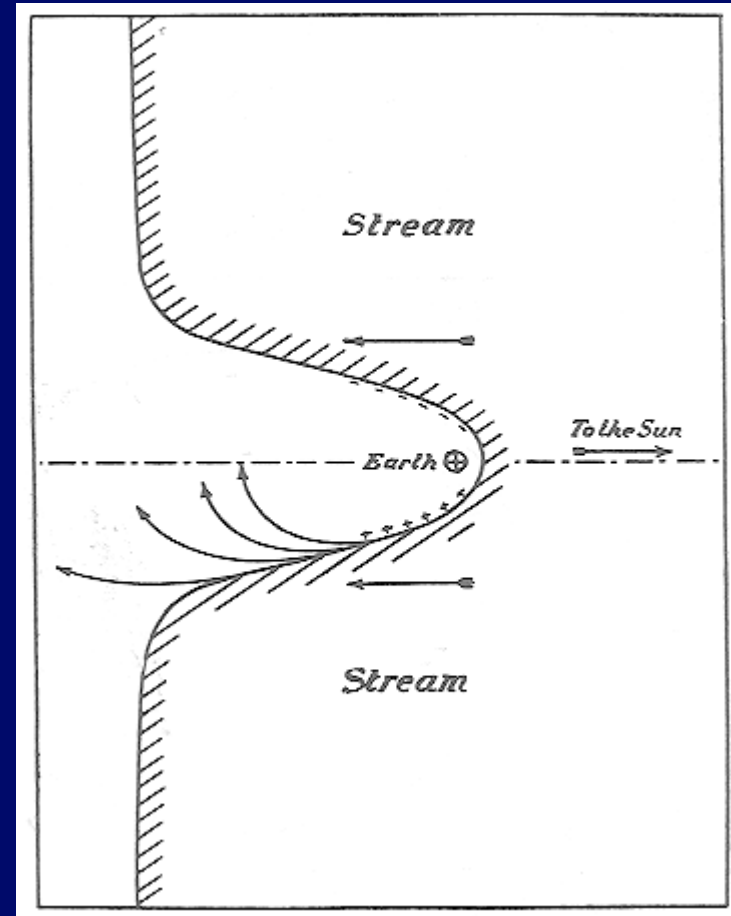
Interaction of the solar wind with the Earth's dipole field

- (1) The Chapman-Ferraro magnetopause
(a closed magnetosphere)

Compression of a dipole field being approached by a rigid superconducting slab



Cavity formation and charge separation through the interaction between a highly conducting medium and a dipole field



Figures from Chapman and Bartels, "Geomagnetism", 1940

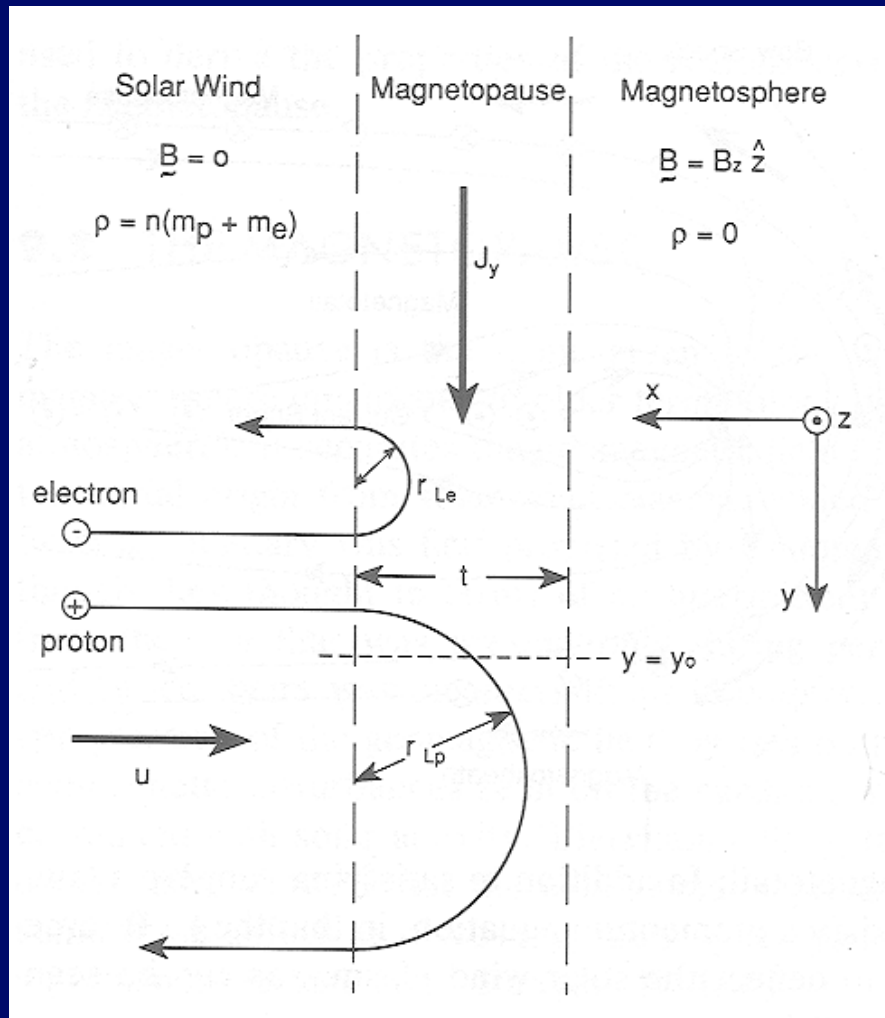


Figure from Hughes, "Magnetopause, magnetotail, reconnection" Chapter 9 in "Introduction to Space Physics"

Larmour radius of protons and electrons

$$r_{Lp,e} = v m_{p,e} / (e B_z)$$

Protons (electrons) crossing $y=y_0$ per time unit are all those which arrive between y_0 and $y_0 - 2 r_{Lp}$ ($y_0 + 2 r_{Le}$)

$$N_{p,e} = 2 r_{Lp,e} n v$$

Current crossing $y=y_0$ per time unit

$$J = (N_p + N_e) e = 2 (r_{Lp} + r_{Le}) n v e$$

$$= 2 \{v (m_p + m_e) / (e B_z)\} n v e$$

For full repulsion the current cancels the magnetic field outside the magnetopause

$$J = B_z / \mu_0$$

$$\rho_{SW} = n (m_p + m_e)$$

Combing all yields the balance between solar wind dynamic pressure and magnetospheric magnetic pressure

$$\rho_{SW} v_{SW}^2 = B_{MS}^2 / (2 \mu_0)$$

solar wind dynamic pressure

$$\rho v^2$$

solar wind magnetic and thermal pressure

$$p$$

momentum conservation

(integration over stream tube)

$$(\rho v^2 + p) S = \text{const}$$

far away from magnetopause

$$p_\infty \ll \rho_\infty v_\infty^2$$

at magnetopause nose (stagnation point)

$$\rho_{\text{stag}} v_{\text{stag}}^2 \ll p_{\text{stag}}$$

Mach number (standard solar wind)

$$M \approx 6$$

after some theory

$$K = p_{\text{stag}} / (\rho_\infty v_\infty^2)$$

$$= \left\{ \frac{1}{2} (\gamma + 1) \right\}^{(\gamma + 1)/(\gamma - 1)} \left\{ \frac{\gamma - (\gamma - 1)}{2M^2} \right\}^{\gamma - 1}$$

$$\approx 0.9 \text{ (typical solar wind conditions)}$$

inside magnetopause

$$p_{\text{plasma}} \ll p_{\text{magnetic}}$$

pressure balance at magnetopause

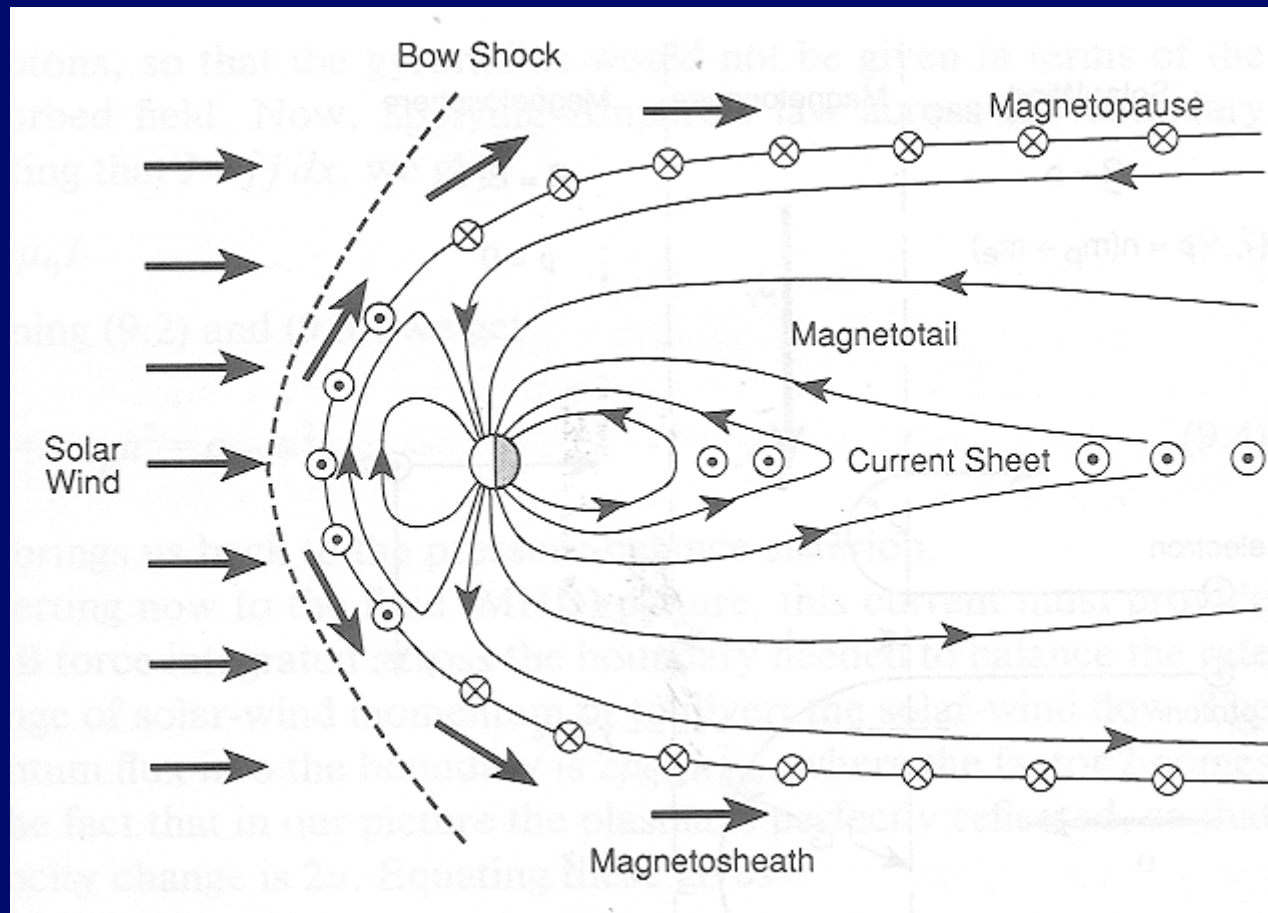
$$K \rho_\infty v_\infty^2 = (\alpha B_0)^2 / (2 \mu_0 L_{\text{mp}}^6)$$

compression factor (accounts for plasma)

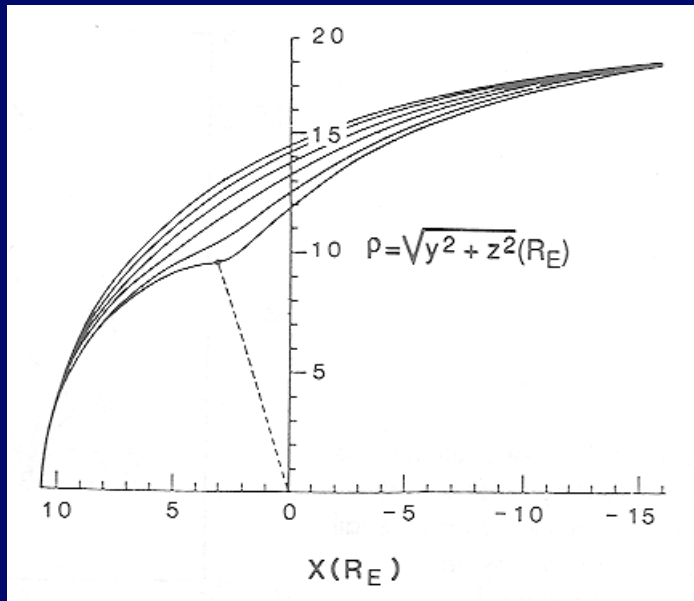
$$\alpha \approx 2-3$$

Interaction of the solar wind with the Earth's dipole field

- (1) The Chapman-Ferraro magnetopause
(a closed magnetosphere)
- (2) The magnetospheric cavity



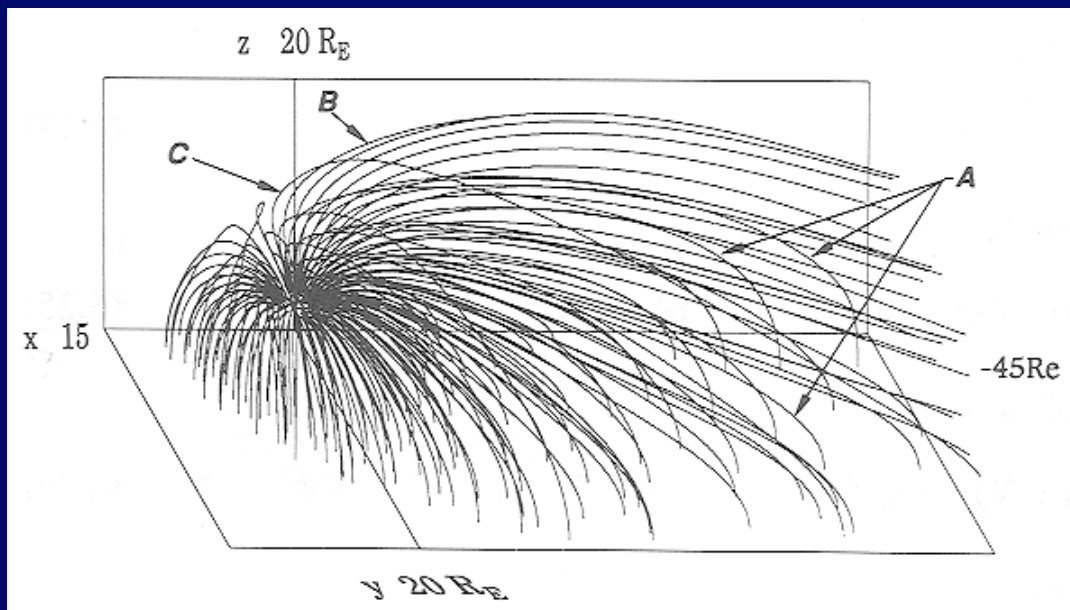
*Figure from Hughes, "Magnetopause, magnetotail, reconnection"
Chapter 9 in "Introduction to Space Physics"*

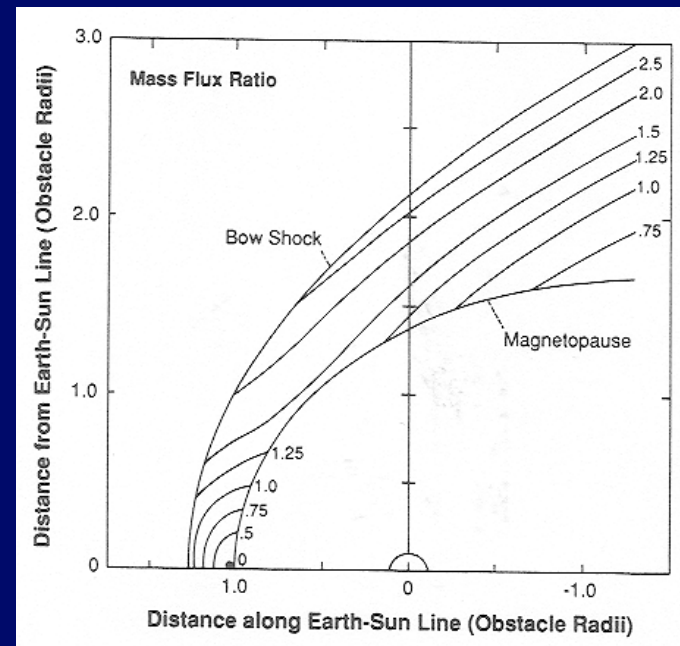
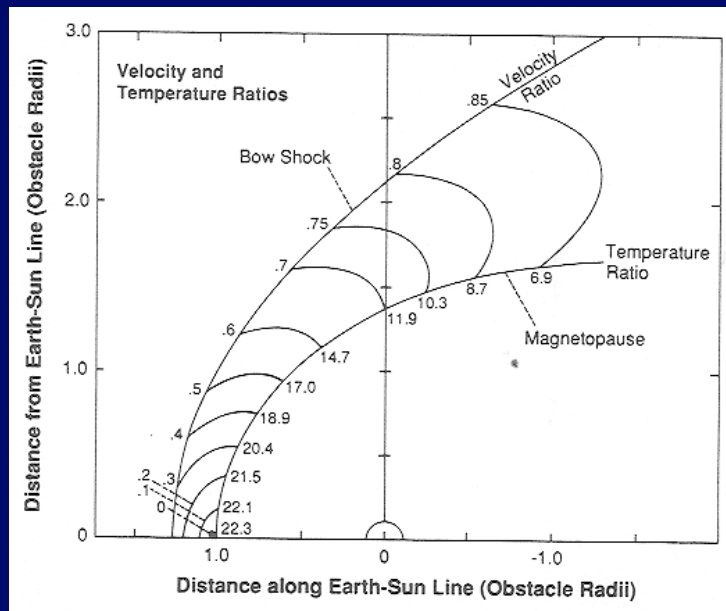
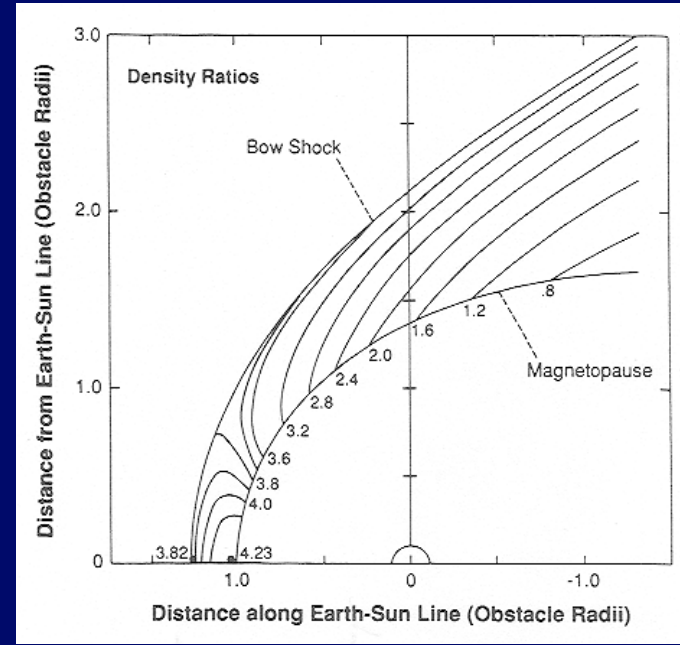
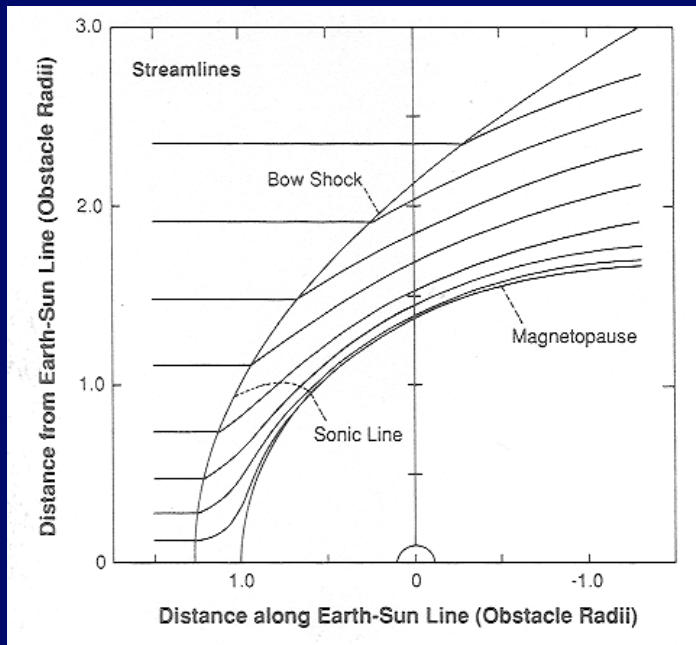


Numerical models of the magnetosphere

Top figure after Mead and Beard (1964)

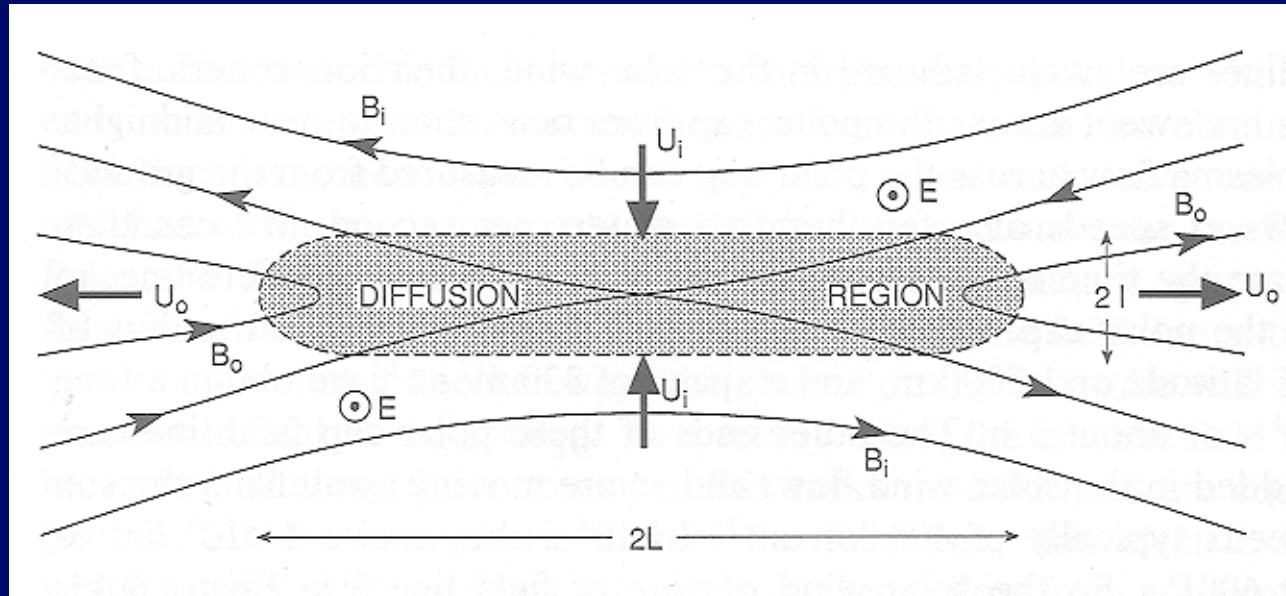
Bottom figure from Ogino et al. (1992)





Figures after Spreiter et al. (1966)

The concept of field line merging
(magnetic reconnection)



Sweet-Parker reconnection geometry

Figure from Hughes, "Magnetopause, magnetotail, reconnection" chapter 9 in "Introduction to Space Physics"

Magnetic Reynolds number

$$R_m = \mu_0 \sigma v L \quad (L = \text{typical scale size of the system})$$

$$R_m \gg 1 \Rightarrow \text{frozen-in flux condition holds}$$

General magnetospheric flow:

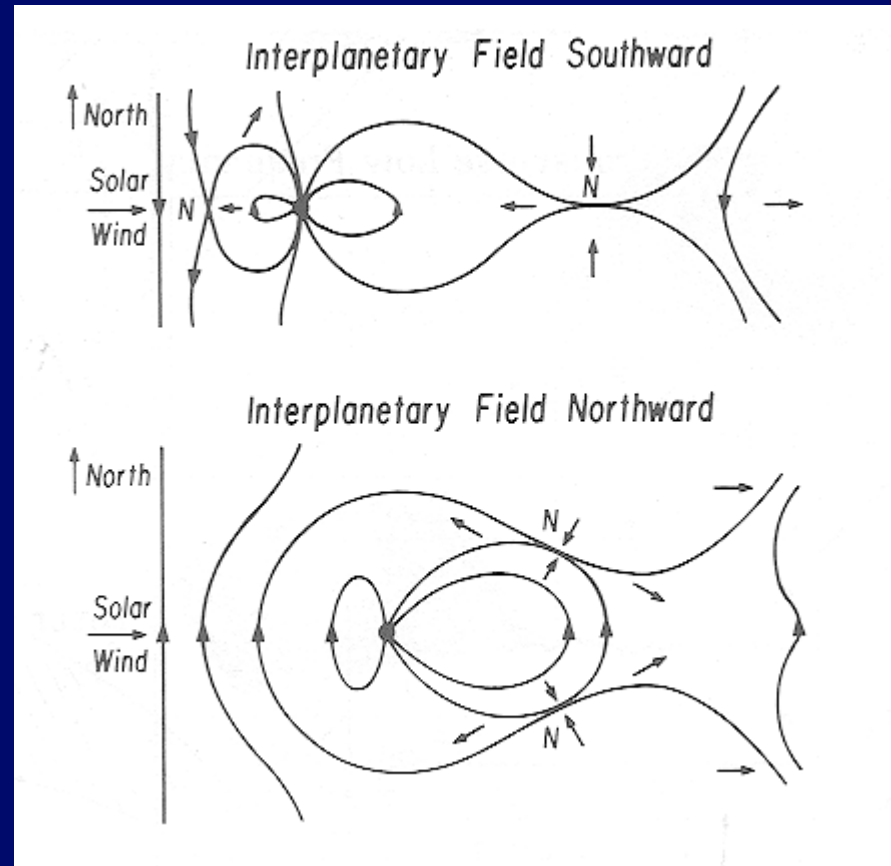
$$v \approx 100 \text{ km/s}, L \approx R_E \Rightarrow R_m \gg 1$$

Neighbouring flux tubes can be significantly different

- they are separated by a thin current sheet
 - L can be small here such that $R_m \approx 1$
 - a plasma diffusion region is generated
 - reconnection can occur at a sufficiently fast rate

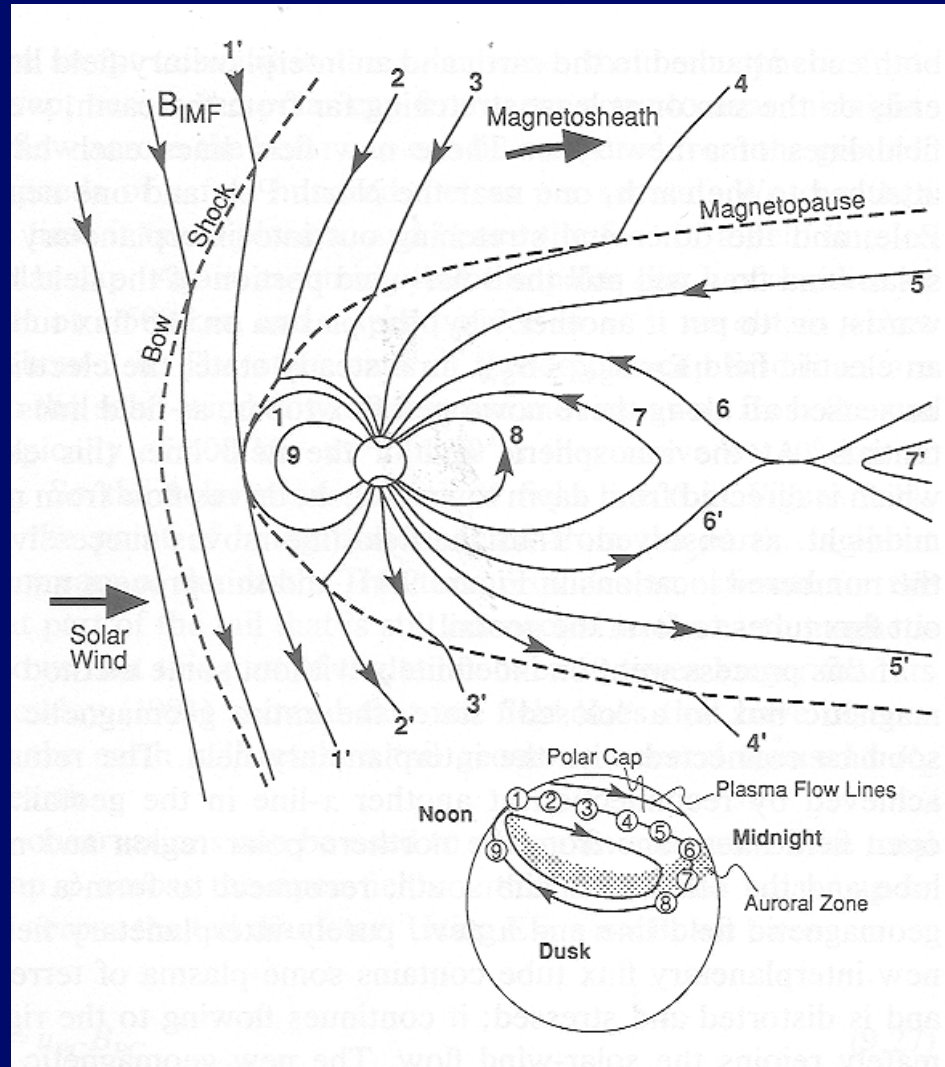
Interaction of the solar wind with the Earth's dipole field

- (1) The Chapman-Ferraro magnetopause
(a closed magnetosphere)
- (2) The magnetospheric cavity
- (3) **Dungey's open magnetosphere**



An early model of the open magnetosphere

Dungey (1963)



Dungey's open magnetosphere model: "field line dragging"

Figure from Hughes, "Magnetopause, magnetotail, reconnection",
Chapter 9 in "Introduction to Space Physics"

Interaction of the solar wind with the Earth's dipole field

- (1) The Chapman-Ferraro magnetopause
(a closed magnetosphere)
- (2) The magnetospheric cavity
- (3) Dungey's open magnetosphere
- (4) The geomagnetic tail

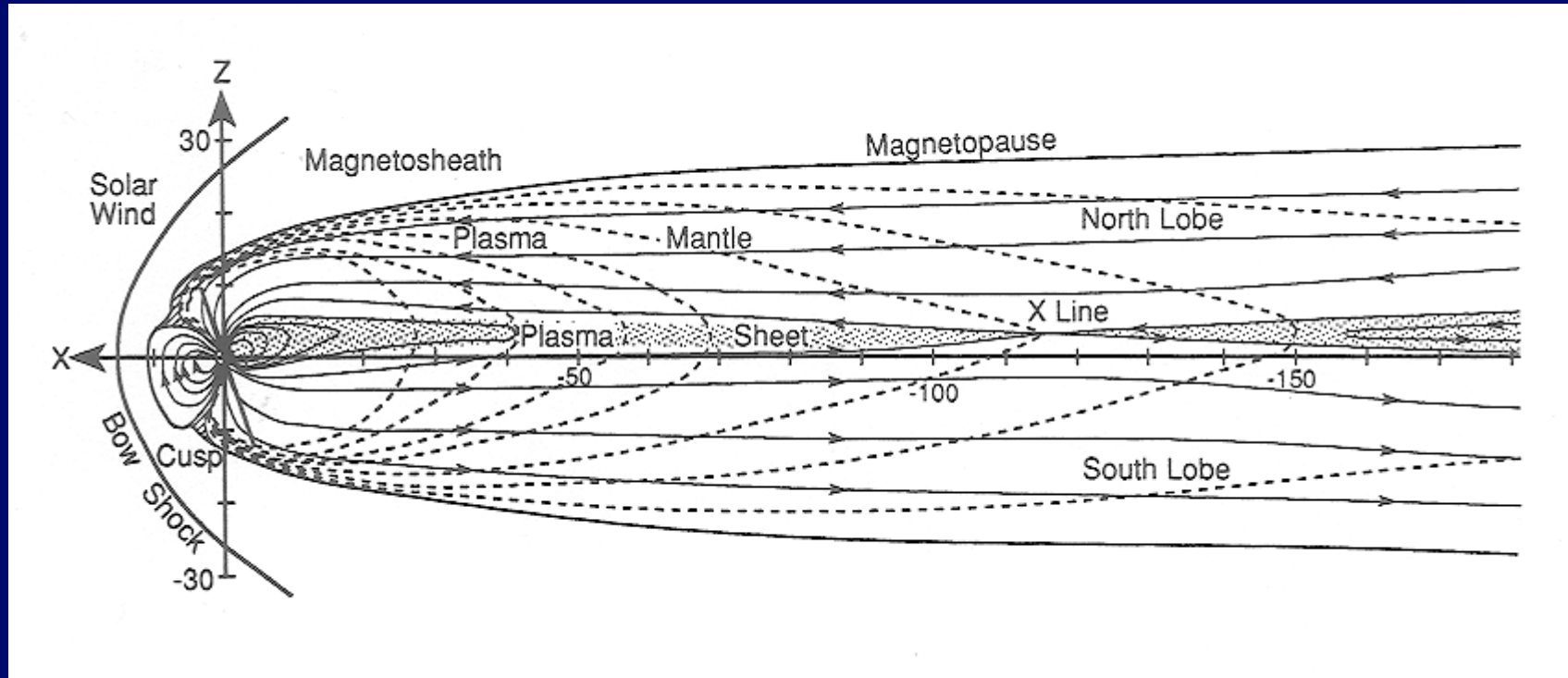
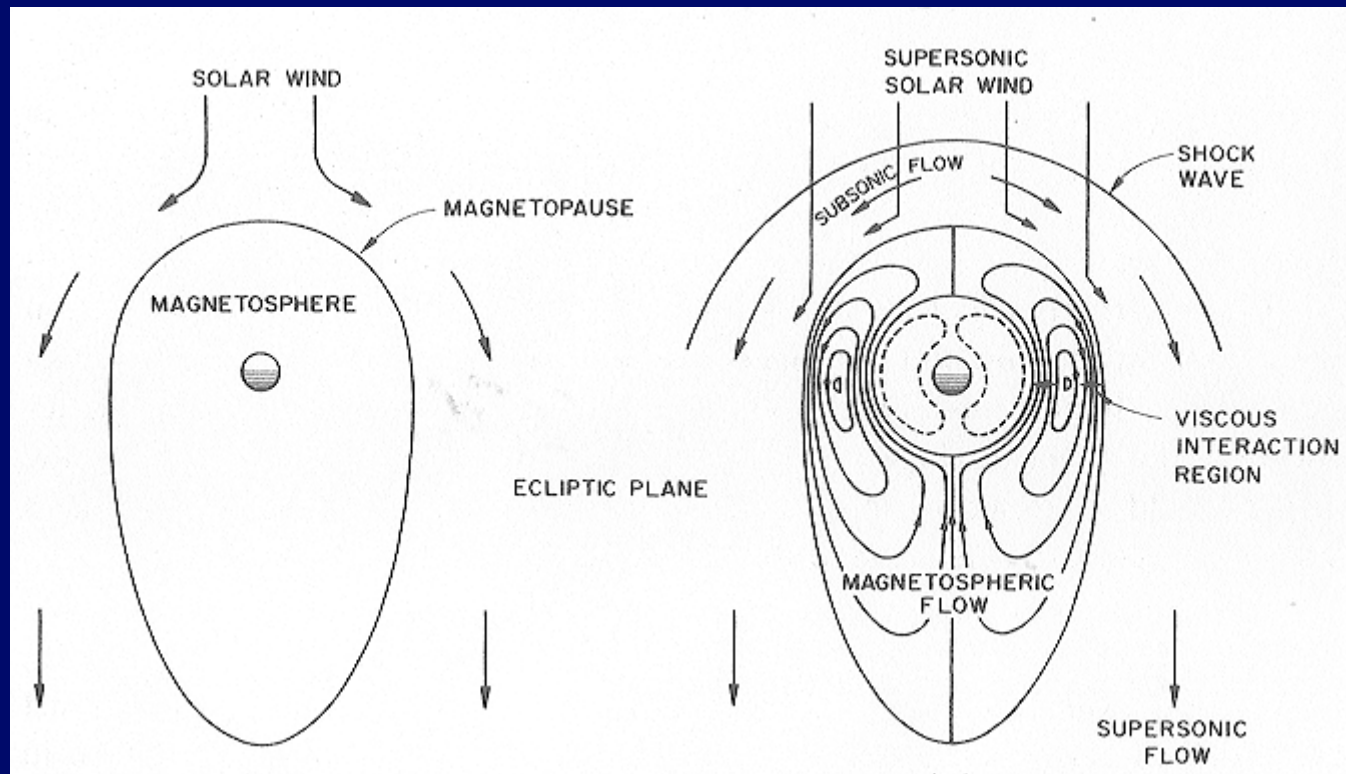


Figure from Hughes, after Philipp and Morfill (1978)

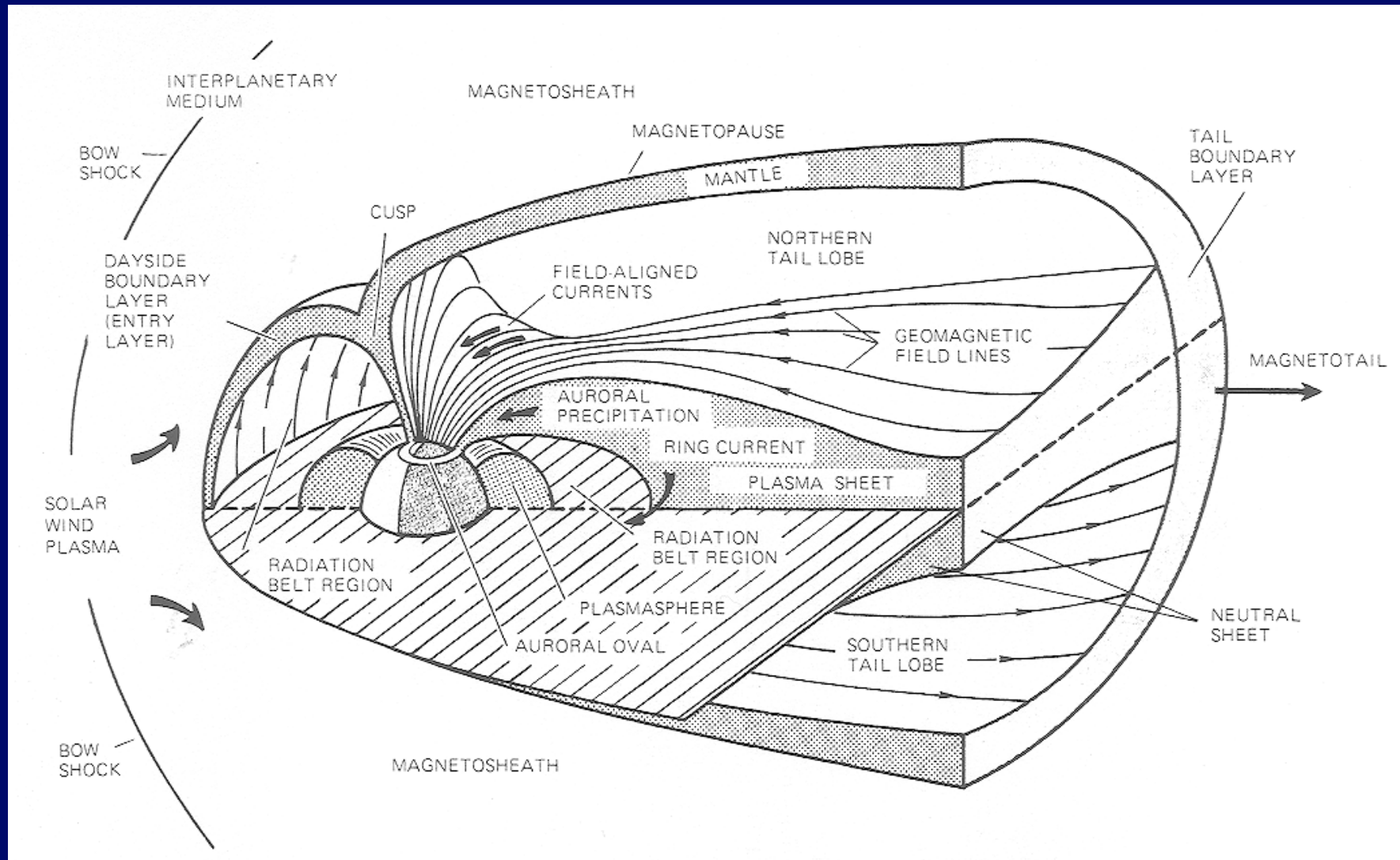
Interaction of the solar wind with the Earth's dipole field

- (1) The Chapman-Ferraro magnetopause
(a closed magnetosphere)
- (2) The magnetospheric cavity
- (3) Dungey's open magnetosphere
- (4) The geomagnetic tail
- (5) Viscous interaction



Viscous interaction across the magnetopause leads to magnetospheric convection. The solar wind is the active part, not the interplanetary magnetic field.

Figure from M.C. Kelley "The Earth's Ionosphere" (1989) after Axford and Hines (1961)



The whole interaction regime: bow shock, magnetosheath, magnetosphere

Figure from M.C. Kelley "The Earth's Ionosphere" (1989)

So far neglected

The bow shock
The magnetosheath

Supersonic shock

airplane: relative air bulk speed exceeds acoustic velocity
mechanical air pressure (collisions) builds up a shock front

solar wind: particle mean free path \gg magnetosphere diameter
 \Rightarrow shock without collisions

upstream ρ, v, B, p

downstream ρ, v, B, p

MHD theory: mass, momentum, energy are conserved
 \Rightarrow several conditions across shock front apply

$\rho v_{\perp} = \text{const}$ (mass stream per unit area)

$\rho v_{\perp}^2 + p + B_{\perp}^2 / (2 \mu_0) = \text{const}$ (shock normal momentum)

$\rho v_{\parallel} - B_{\parallel} / \mu_0 = \text{const}$ (shock parallel momentum)

$B_{\perp} = \text{const}$ (e.m. boundary condition)

$v_{\parallel} B_{\perp} = v_{\perp} B_{\parallel}$ (e.m. boundary condition)

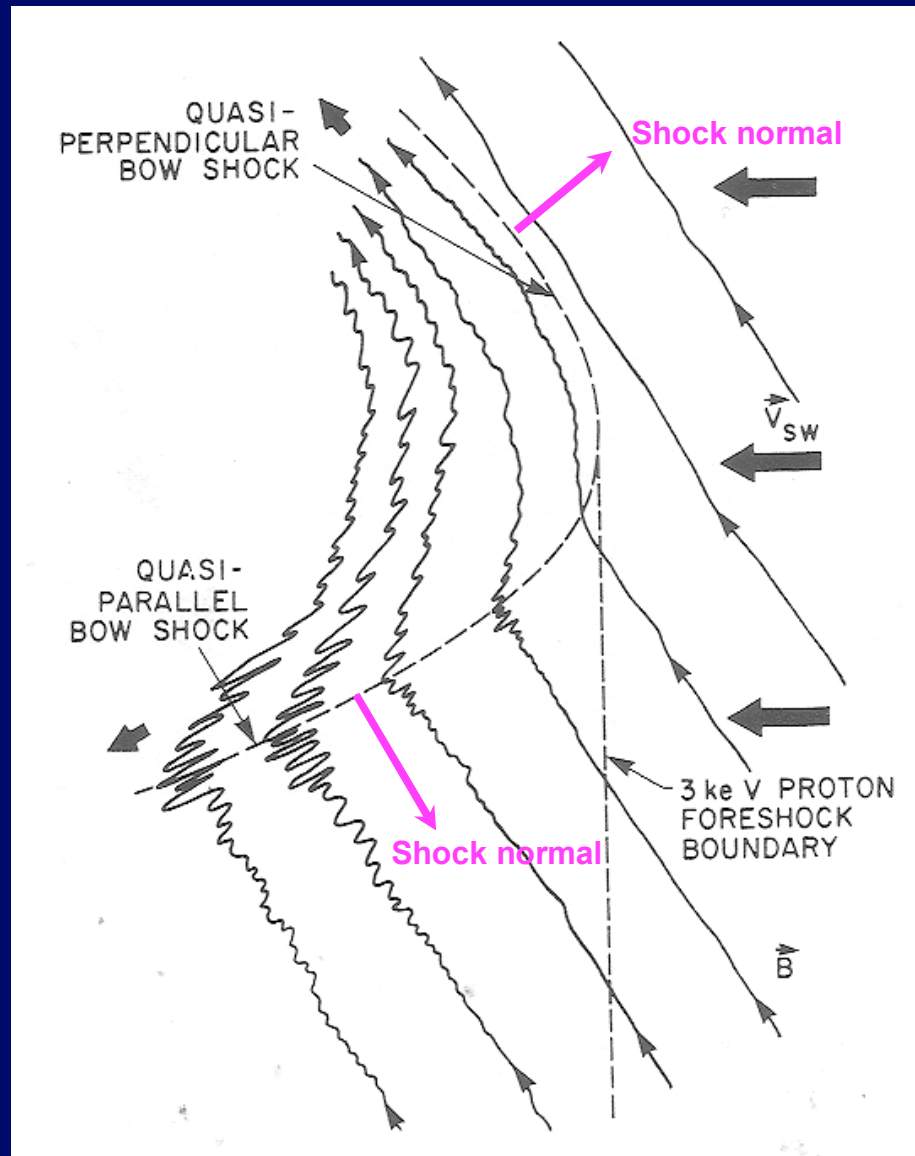


Figure from J. Feynman "Solar Wind"
Chapter 3 in "Handbook of Geophysics and the Space Environment"

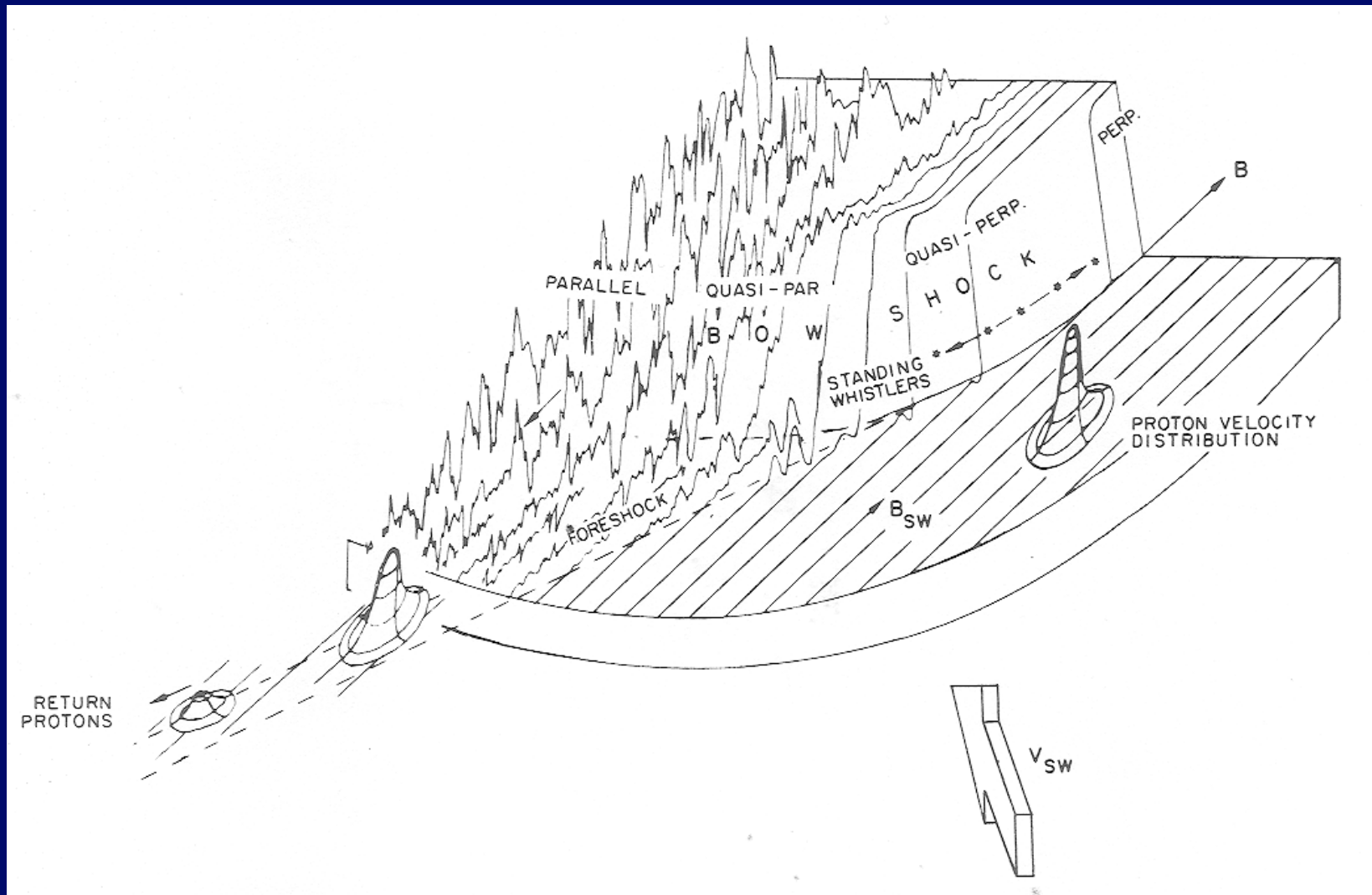


Figure from J. Feynman "Solar Wind"
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MHD model of magnetospheric magnetic field lines and electric currents

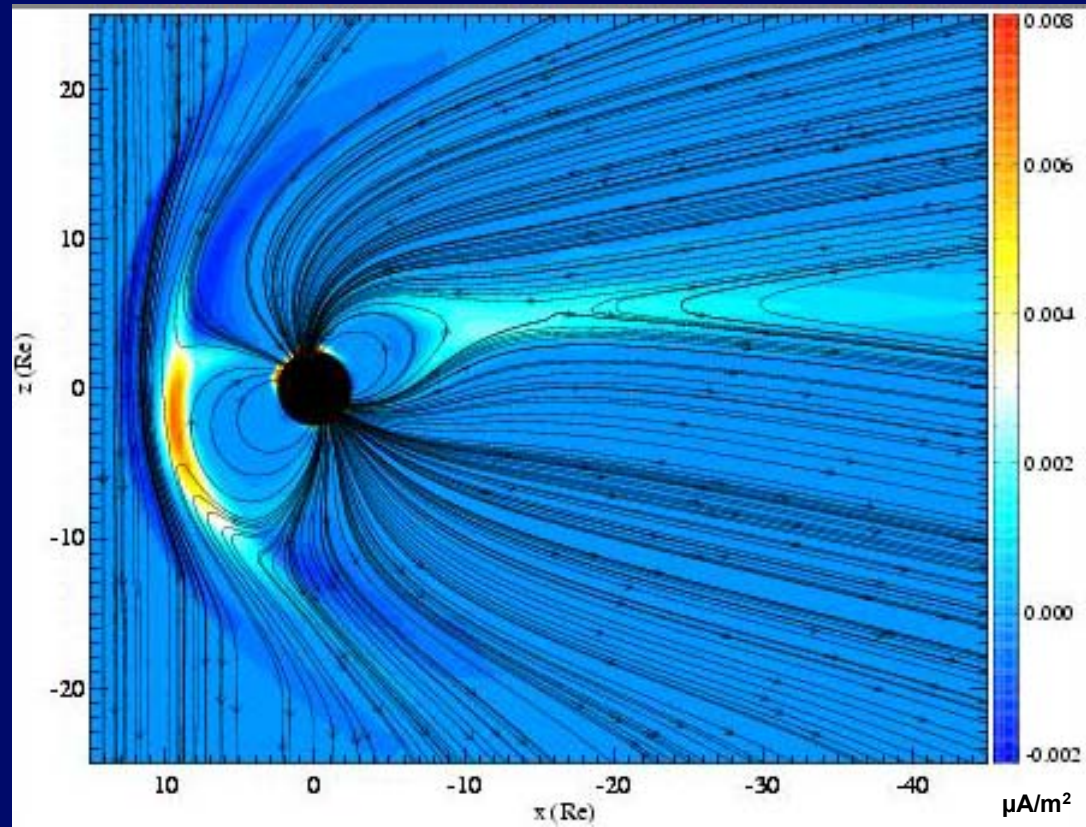
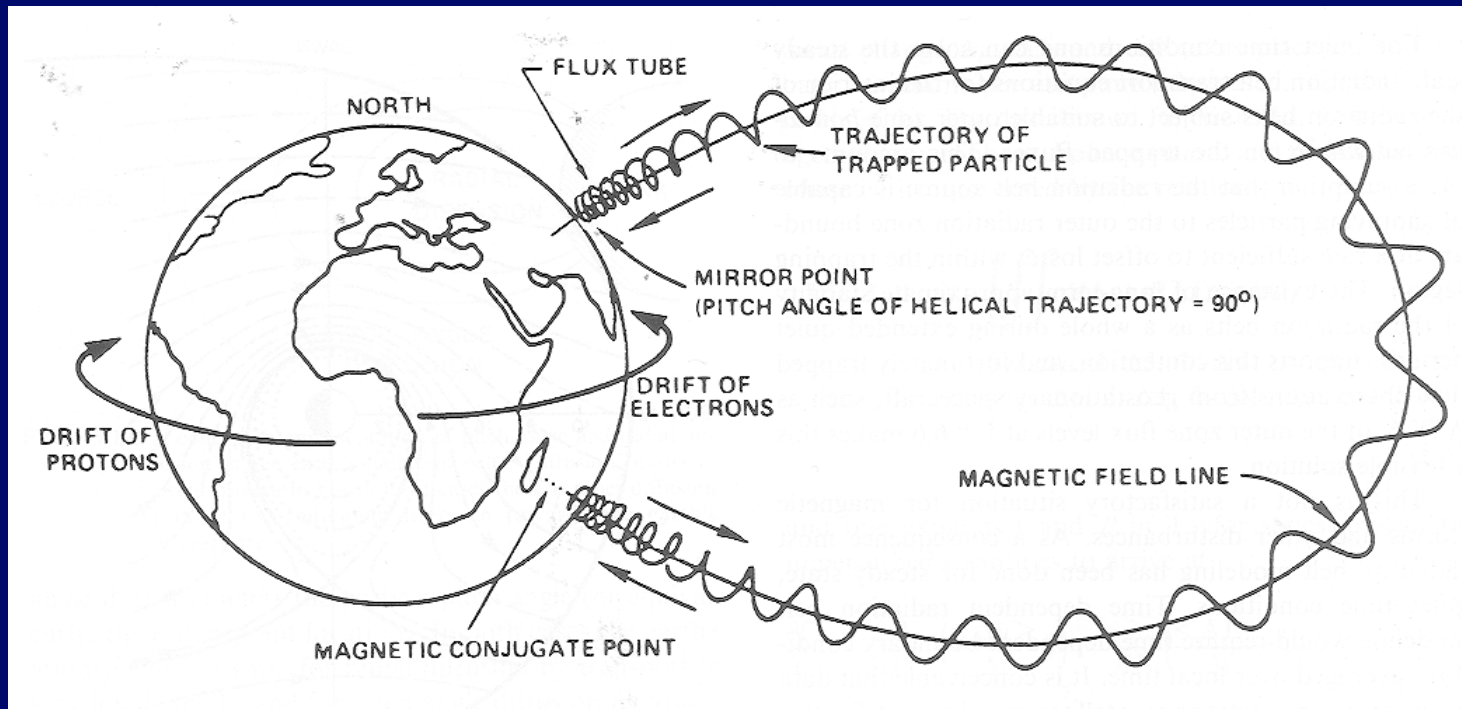


Figure by the Finnish Meteorological Institute (?)

Charged particle motion under stationary conditions

3 adiabatic invariants

(quasi-periodic single particle motions)



quasi-periodic motions at three different time scales

$$f_{\text{gyro}} \gg f_{\text{bounce}} \gg f_{\text{drift}}$$

*Figure from Spjeldvik and Rothwell, "The radiation belts"
Chapter 5 in "Handbook of Geophysics and the Space Environment"*

Hamilton-Jacobi action integral is constant in periodic motions

$$J = \oint p dq = \text{const}$$

Magnetic moment of a particle (first adiabatic invariant)

$$J = \oint mv_{\perp} r_L d\theta = 2\pi r_L m v_{\perp} = 2\pi m v_{\perp}^2 / \omega_c = 4\pi m / |q| \mu$$

Conditions: $B / |\nabla B| \gg mv_{\perp} / (|q|B)$ (*spatial gradient scale* \gg *gyro radius*)

$T_B \gg 2\pi m / (|q|B)$ (*temporal change scale* \gg *gyro period*)

Half-cycle bouncing period between mirror points (second adiabatic invariant)

$$J = \oint v_{\parallel} ds$$

Drift period around the globe (third adiabatic invariant)

$$J = q \oint \underline{B} d\underline{S} = -2\pi B_E R_E^2 / L$$

$L =$ McIlwain parameter, defined by

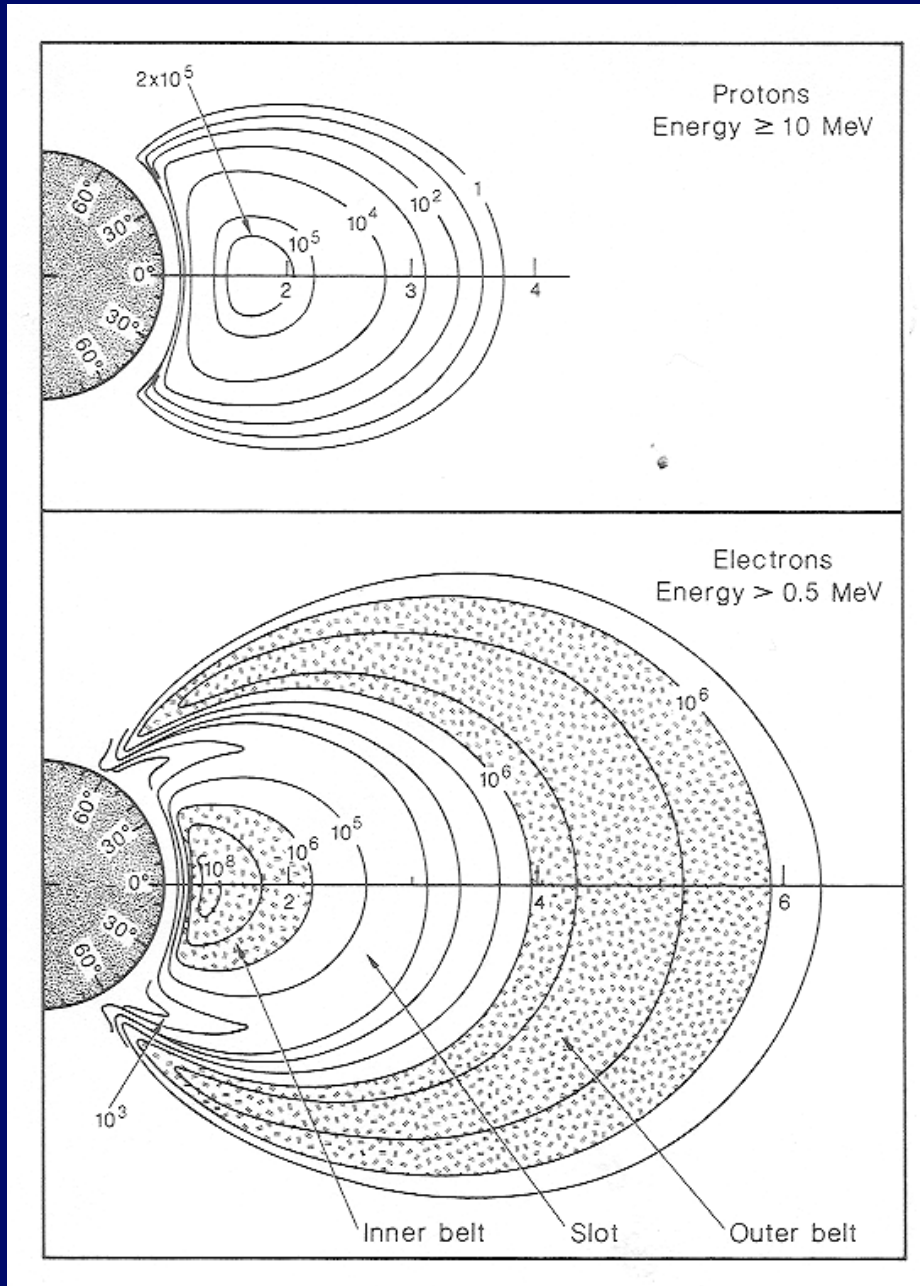
$$L = 1 / \cos^2(\Lambda)$$

with Λ being the latitude where the field line intersects the Earth's surface

Net result

Particles can be trapped in their flux tube thereby drifting around the Earth

- ⇒ high energy particles (several MeV) populate the Inner Van Allan Belts (also termed trapped radiation belts)
- ⇒ high energy particles (several MeV) populate the outer radiation belts
- ⇒ low energy particles (up to several hundred keV) populate the ring current (coincides with outer radiation belt)

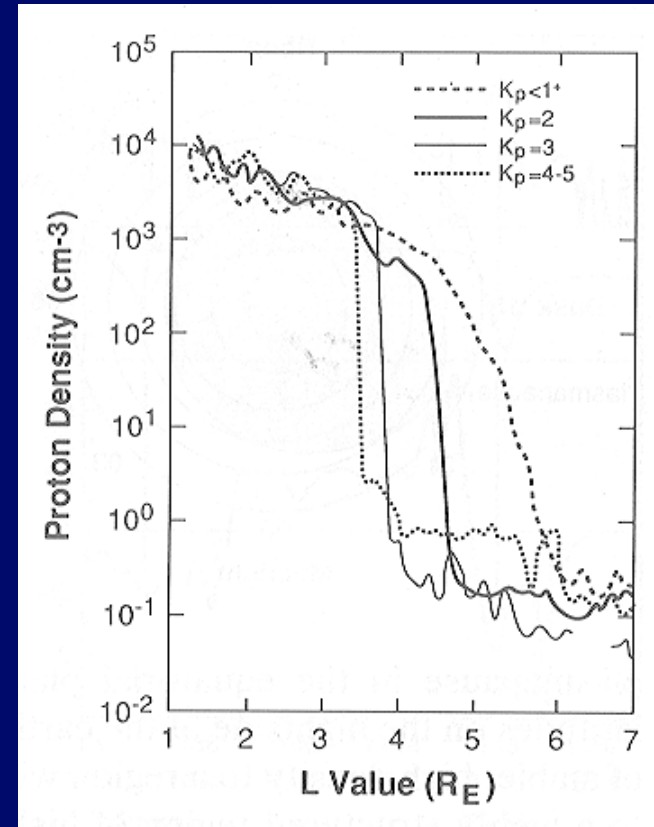
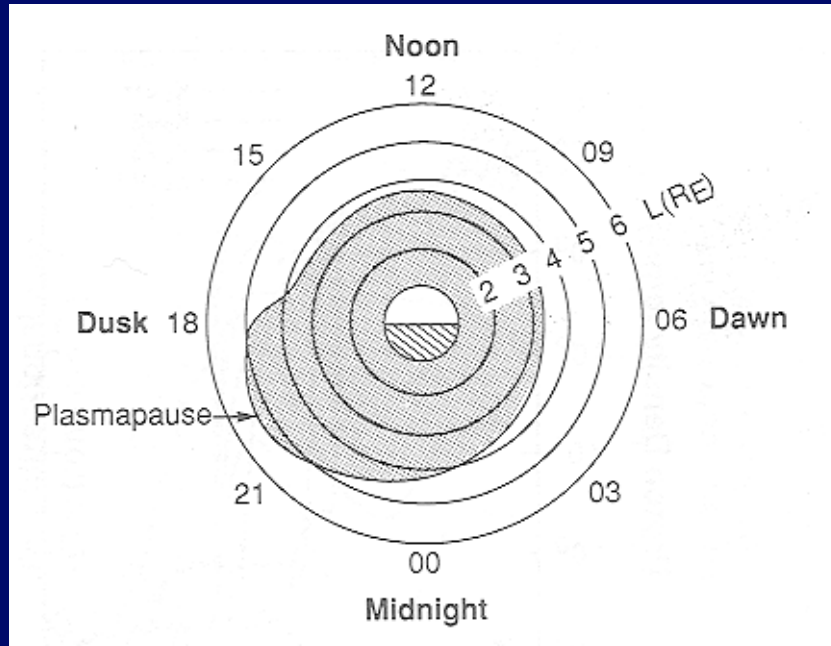


Ion and electron radiation belts

*Figure from Russel, "A brief history of solar-terrestrial physics"
Chapter 1 in "Introduction to space physics"*

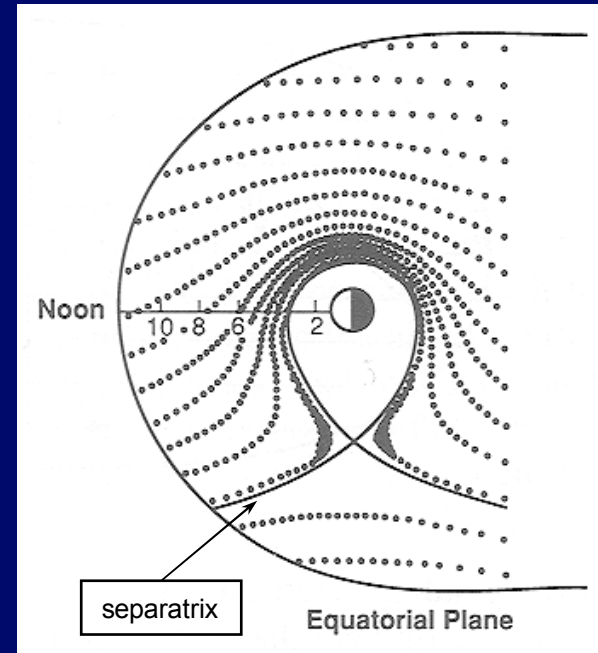
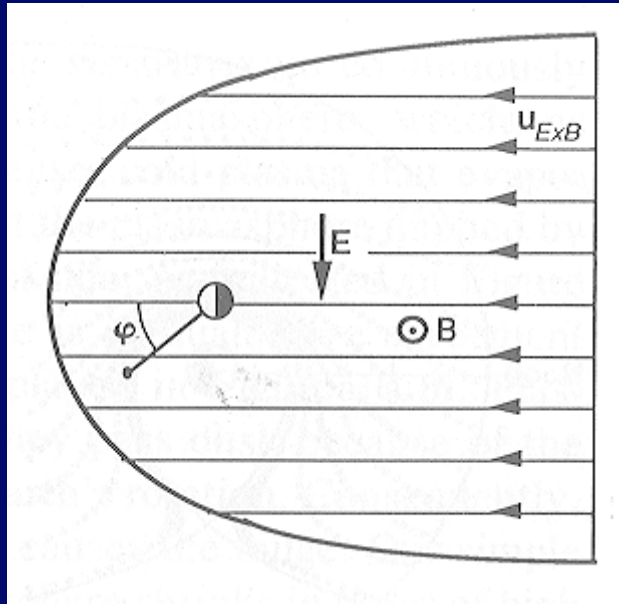
The plasmasphere

Cold (1 eV) plasma in the inner magnetosphere
coexisting with the hot radiation belt population



Plasma density in the inner magnetosphere

*Figures from R. Wolf "Magnetospheric Configuration"
Chapter 10 in "Introduction to Space Physics"*



Figures from R.A. Wolf "Magnetospheric configuration"
Chapter 10 in "Introduction to Space Physics"

Origin of the plasmasphere:

Far-out plasma drift governed by

Near-Earth plasma tends to corotate

Dawn side of the Earth

Dusk side of the Earth

outflow of plasma from polar regions

$$\Phi_{\text{convection}} = - E r \sin\phi$$

$$\Phi_{\text{corotation}} = - \omega B_0 r_E^3 / r$$

both drive plasma sunward

they counteract each other

Separatrix forms a strict boundary which cannot be crossed

Magnetospheric current system

(magnetosphere-ionosphere coupling currents)

(1) The large-scale current system

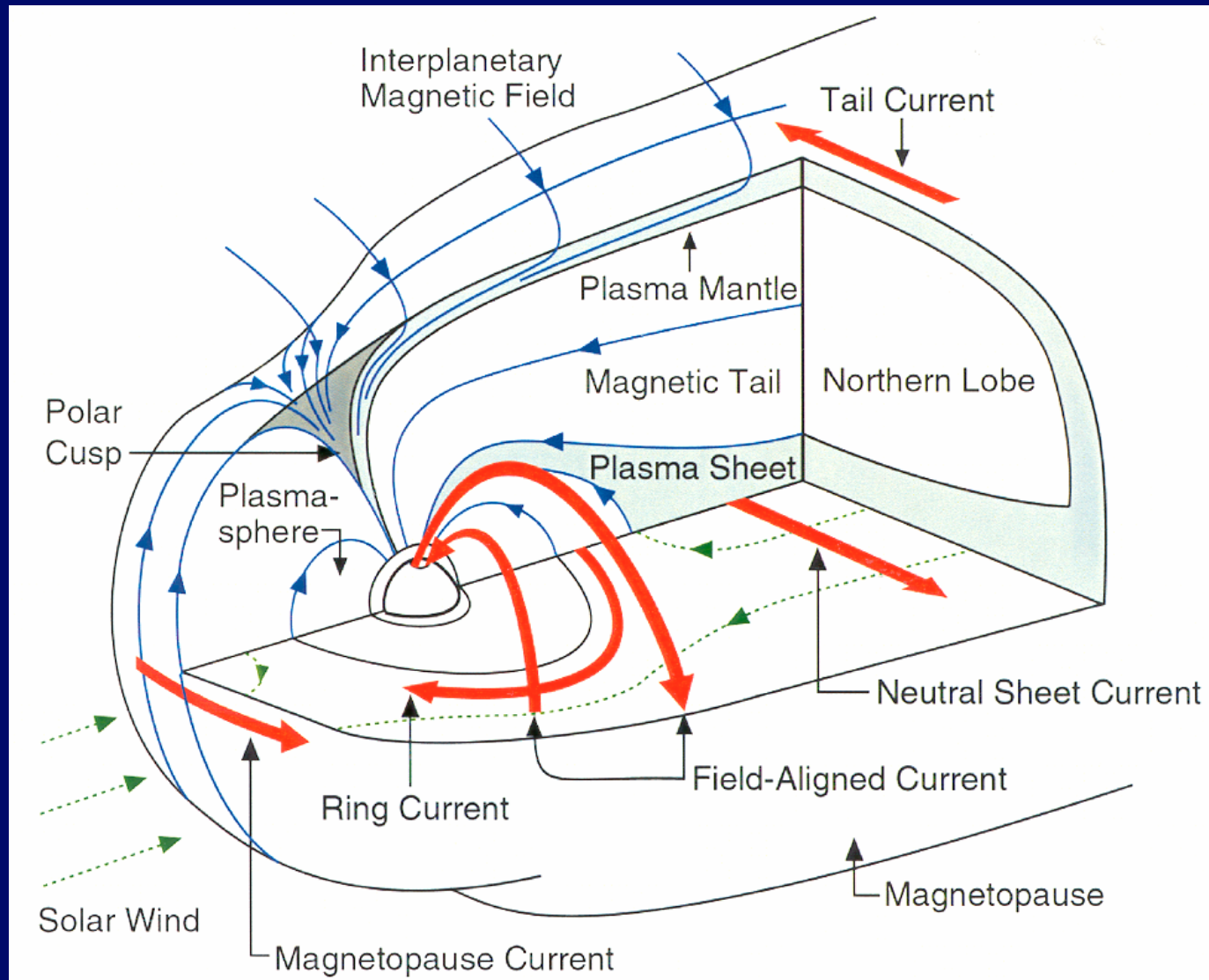
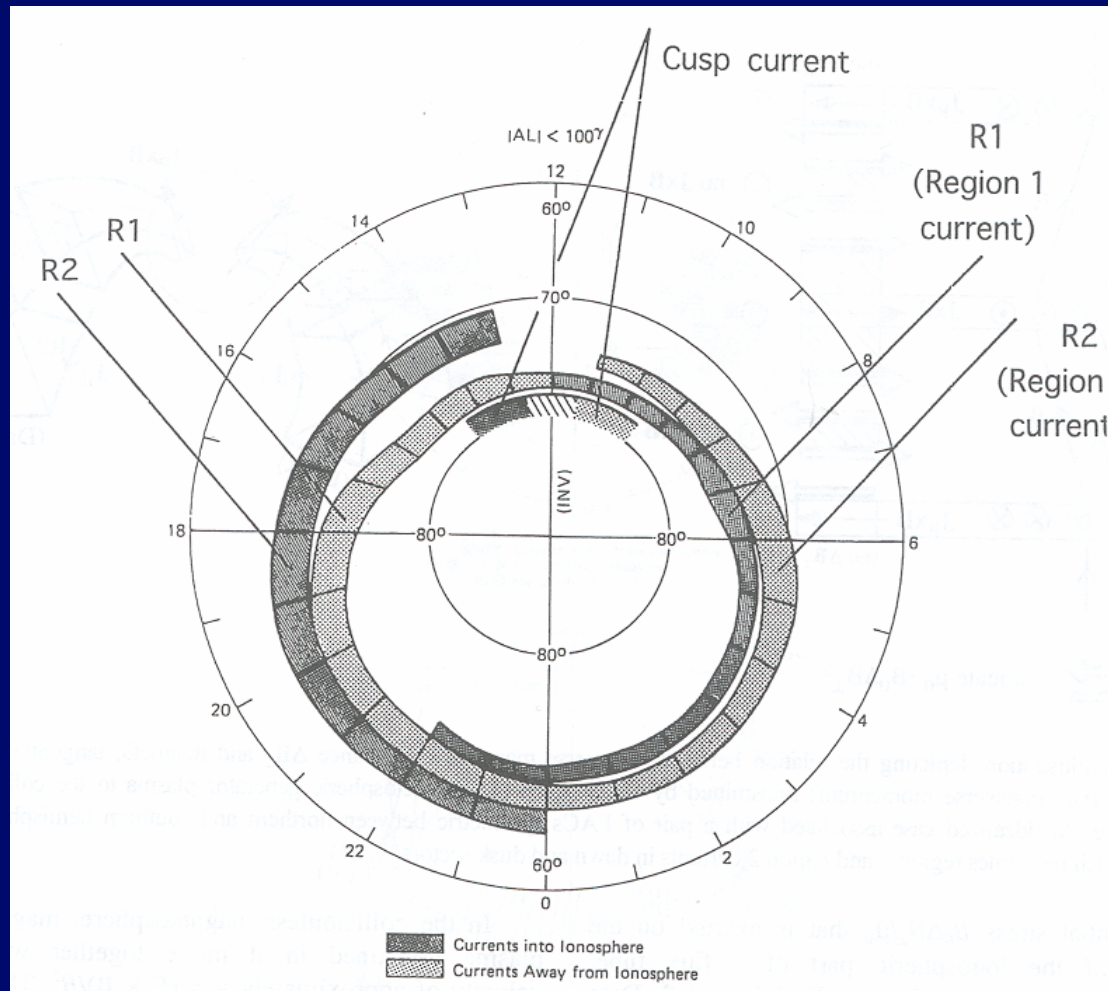


Figure from "Report from the first Geospace Environment Modeling (GEM) Campaigns: 1991-1997"

Magnetospheric current system

(magnetosphere-ionosphere coupling currents)

- (1) The large-scale current system
- (2) Field-aligned currents (FAC)



Basic system of field-aligned currents
 derived from TRIAD satellite measurements

Iijima and Potemra (1982)

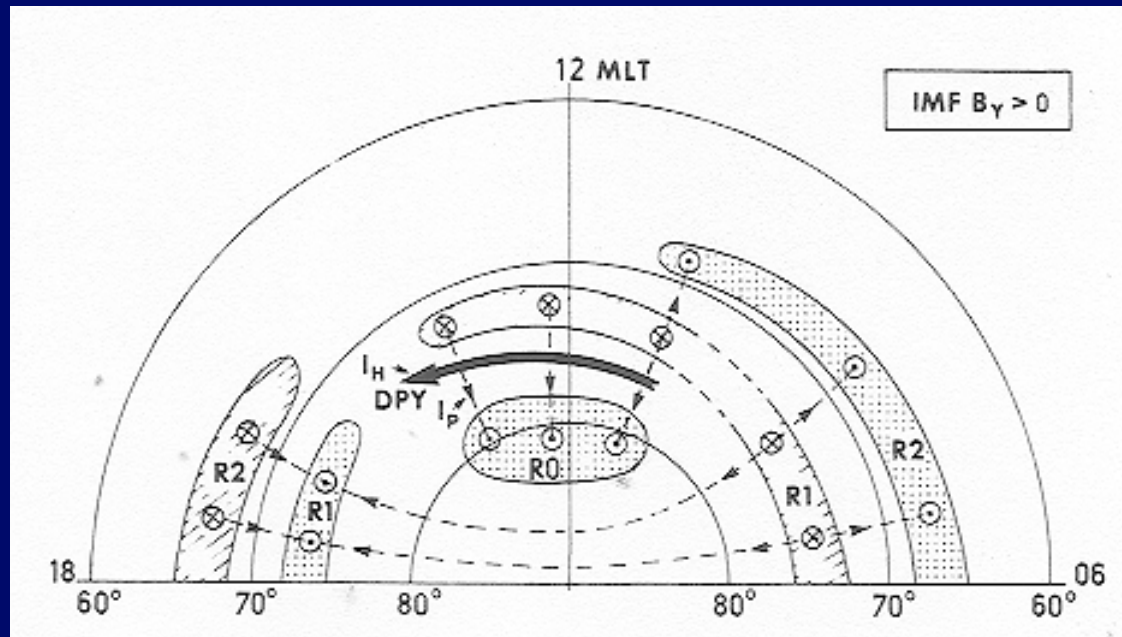
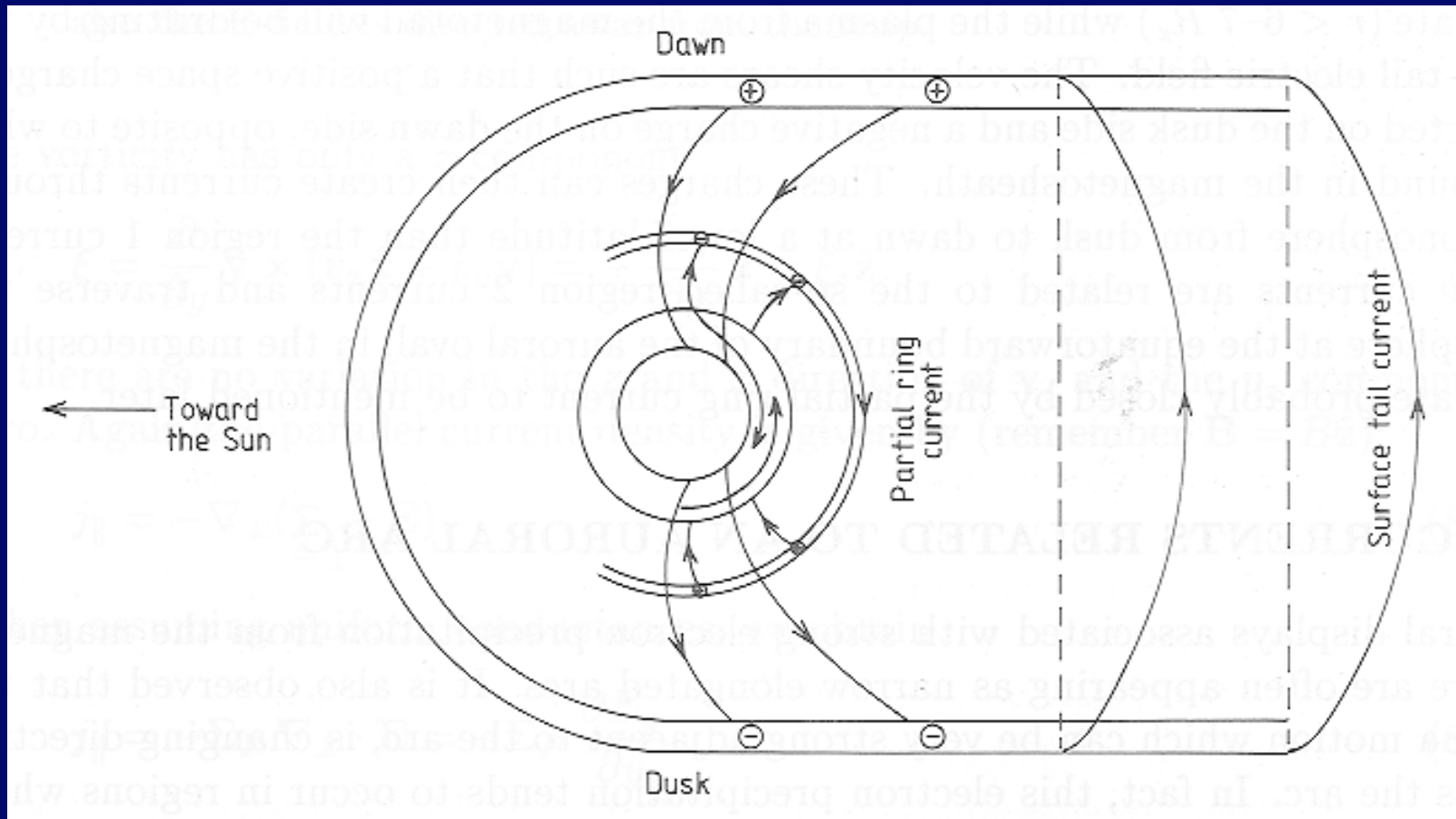


Figure and list by Stauning, DMI

Region 1 (R1) FACs

connect to Low-latitude Boundary Layer (LLBL)

- Region 2 (R2) FACs
connect to Ring Current (RC) Region
- Region 0 (R0) or "Cusp Currents"
connect to Cusp Region (Ionospheric DPY currents)
- Northward IMF B_z (NBZ) FACs
connect to High-latitude Boundary Layer (HLBL)
- Substorm FACs
connect to Tail Region (Plasma Sheet)

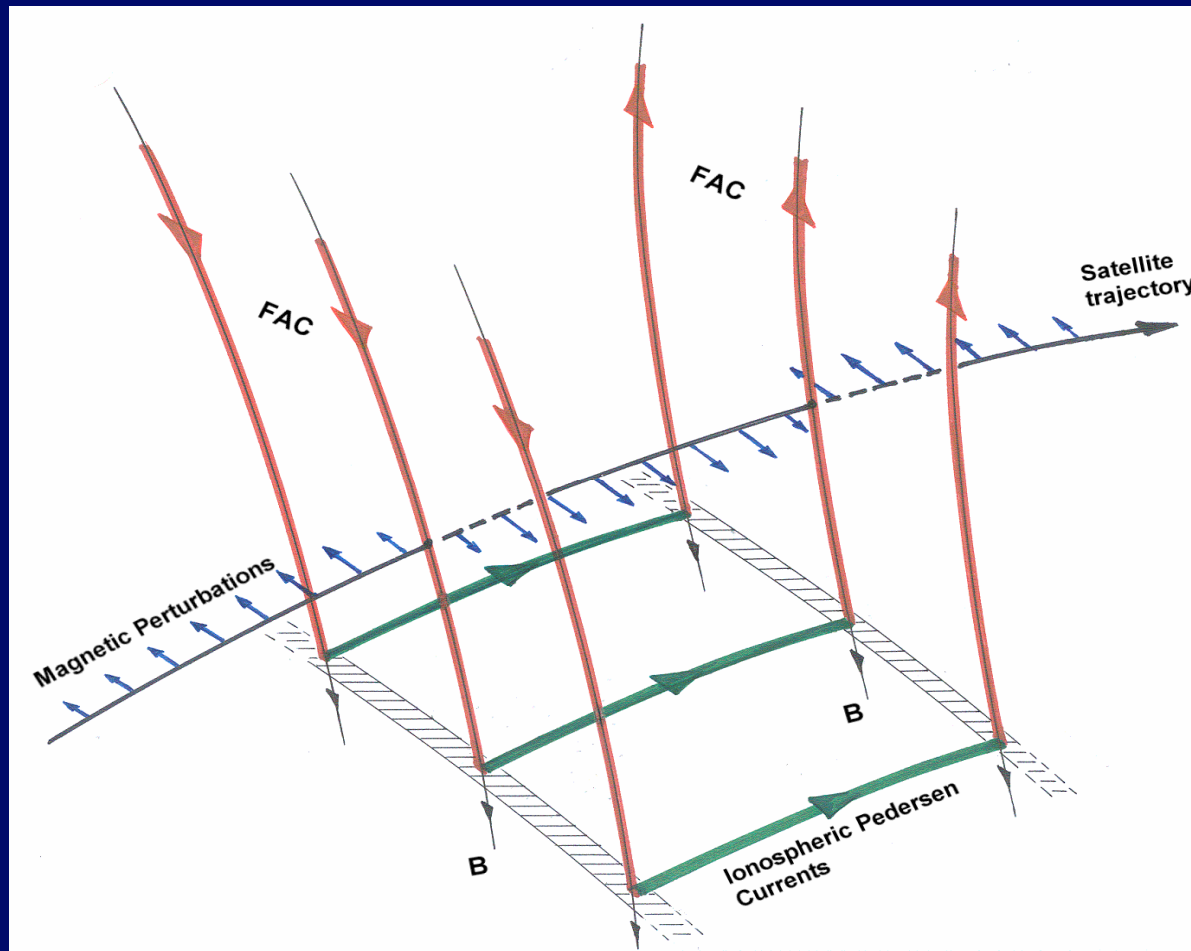


Closure of near-Earth magnetospheric current system

Figure from Brekke "Physics of the Upper Atmosphere" (1997)

How to determine FAC from satellite magnetometer measurements?

(a) Inference from a single satellite pass



Low-altitude satellite pass crossing field-aligned current sheets

Figure prepared by P. Stauning, DMI

FAC inferred from a single satellite pass

Right-handed orthogonal system:

x along satellite trajectory

y perpendicular to satellite trajectory

z along field line (south → north)

$$\Rightarrow B_z = 0$$

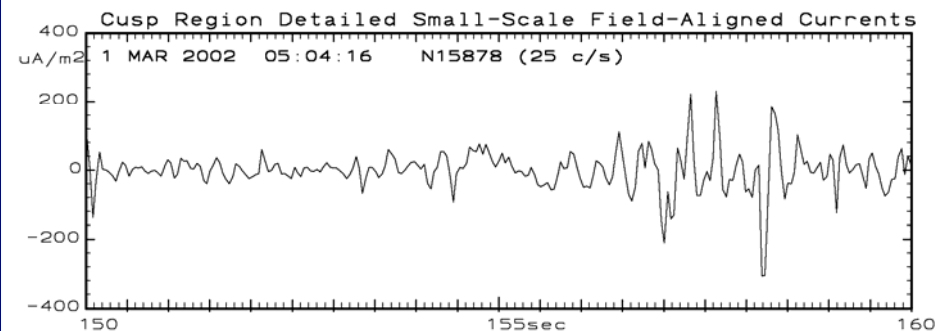
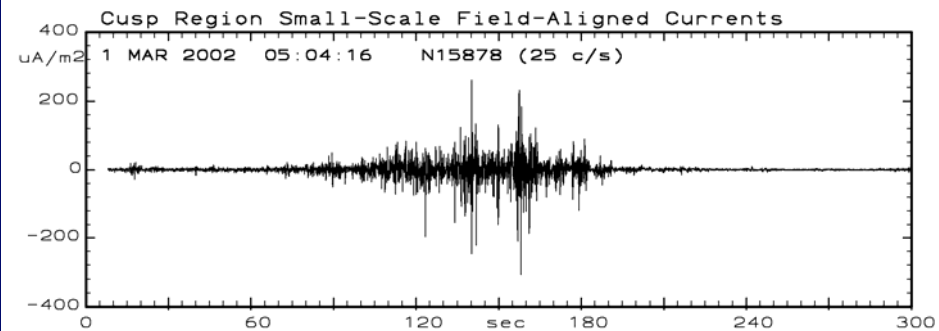
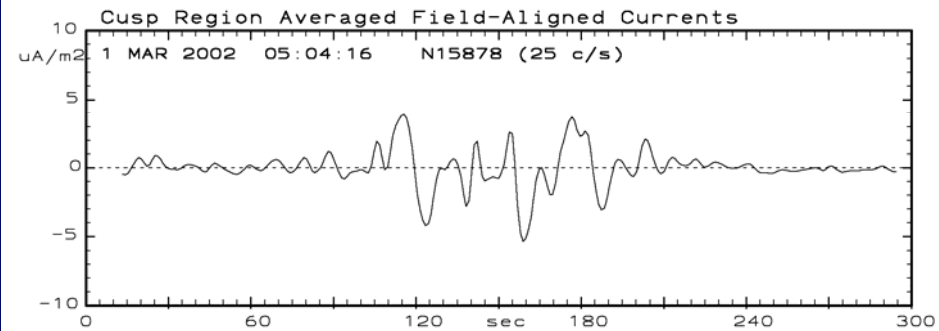
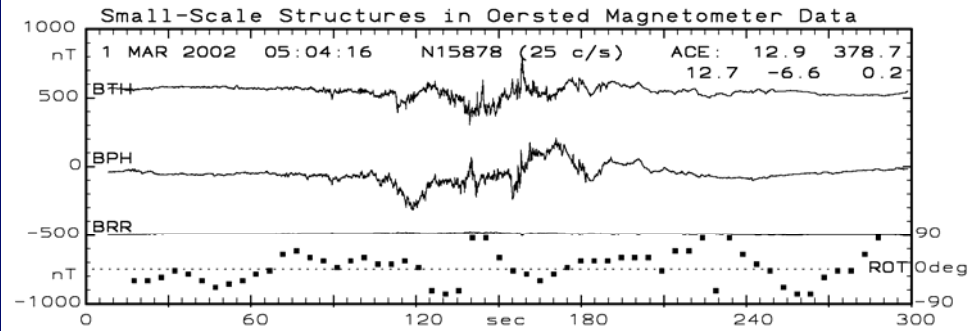
Ampère's law:

$$\mu_0 j_{\parallel} = \nabla \times B = \partial B_y / \partial x - \partial B_x / \partial y$$

Assume FAC in parallel sheets

φ = angle between trajectory and FAC sheets

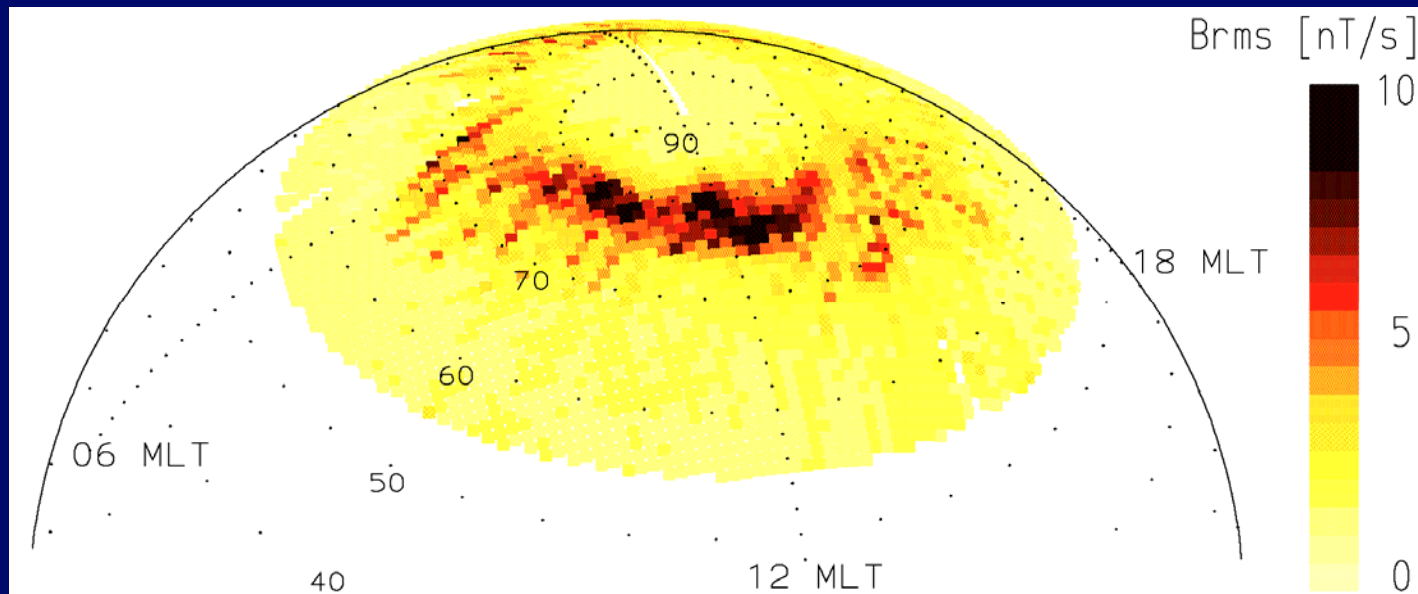
$$\mu_0 j_{\parallel} = 1/\sin\varphi \partial B_y / \partial x \quad (\text{discard data if } \varphi < 45^\circ)$$



Large-scale and small-scale field-aligned currents

Figure by P. Stauning, DMI

Occurrence of fine-scale field-aligned currents



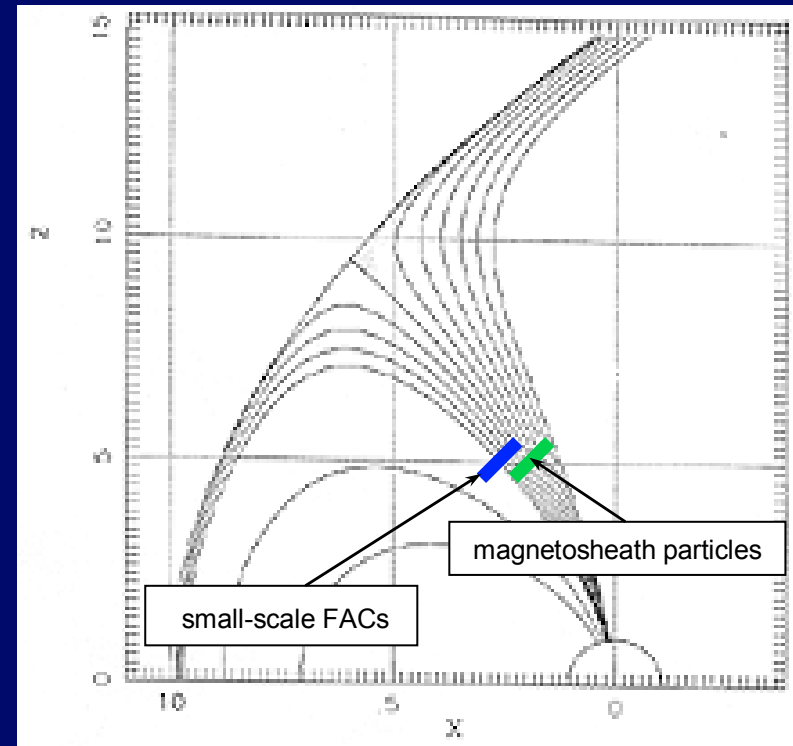
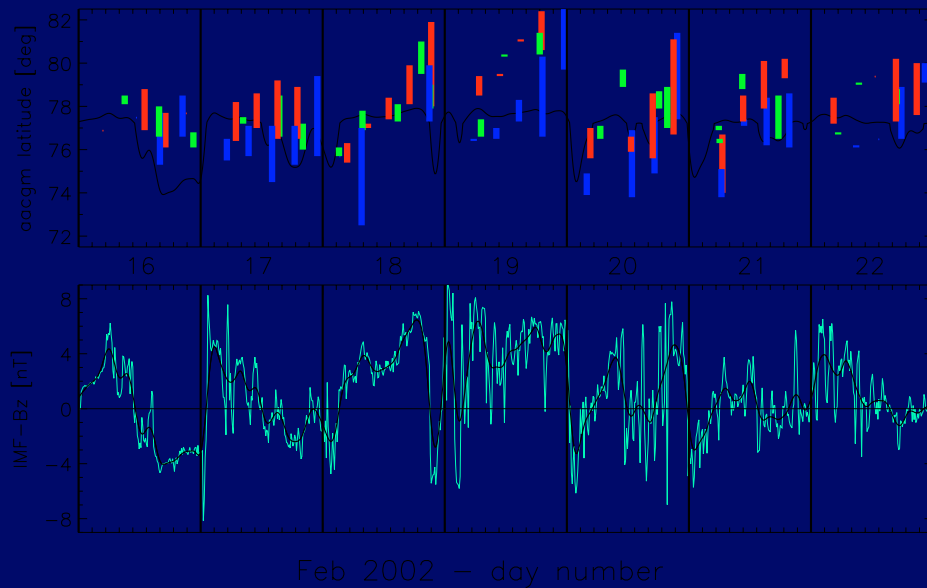
Distribution of fine-scale FAC intensities over the polar cap mapped in magnetic latitude – local time coordinates

Figure from Stauning et al., "Earth Observations with CHAMP", Springer, Berlin - Heidelberg - New York, 2005

Results from a DMSP – Ørsted – CHAMP joint data analysis

LEO satellite observations vs cusp statistical equatorward boundary

DMSP (charged particles) Ørsted (small scale δB) CHAMP (small scale FAC)



How to determine FAC from satellite magnetometer measurements?

- (a) Inference from a single satellite pass
- (b) Spatial interpolation of data accumulated over many orbits

FAC inferred from a multitude of satellite passes

Right-handed orthogonal system:

- x along satellite trajectory
- y perpendicular to satellite trajectory
- z along field line (south → north)

⇒ $B_z = 0$

Ampère's law: $\mu_0 j_{\parallel} = \nabla \times \mathbf{B} = \partial B_y / \partial x - \partial B_x / \partial y$

Bin data from many satellite passes according to certain criteria in order to fill a latitude-longitude grid (*in the hope that very similar magnetospheric conditions prevailed during the observations for each bin*)

- solar wind mass density and velocity
- IMF magnitude and orientation (*determines field line merging*)
- season (*influences critically ionospheric conductivity*)

Interpolate between grid points (e.g., fit 2nd order polynomial to each cell)

Apply Ampère's law on the two-dimensional grid (now full curl)

Steady-state FAC Distributions ($V_{sw} \sim 400$, $D_{sw} \sim 5$, summer conditions)

Figure by P. Stauning, DMI

red : outward
blue : inward

$B_y = -10$ nT

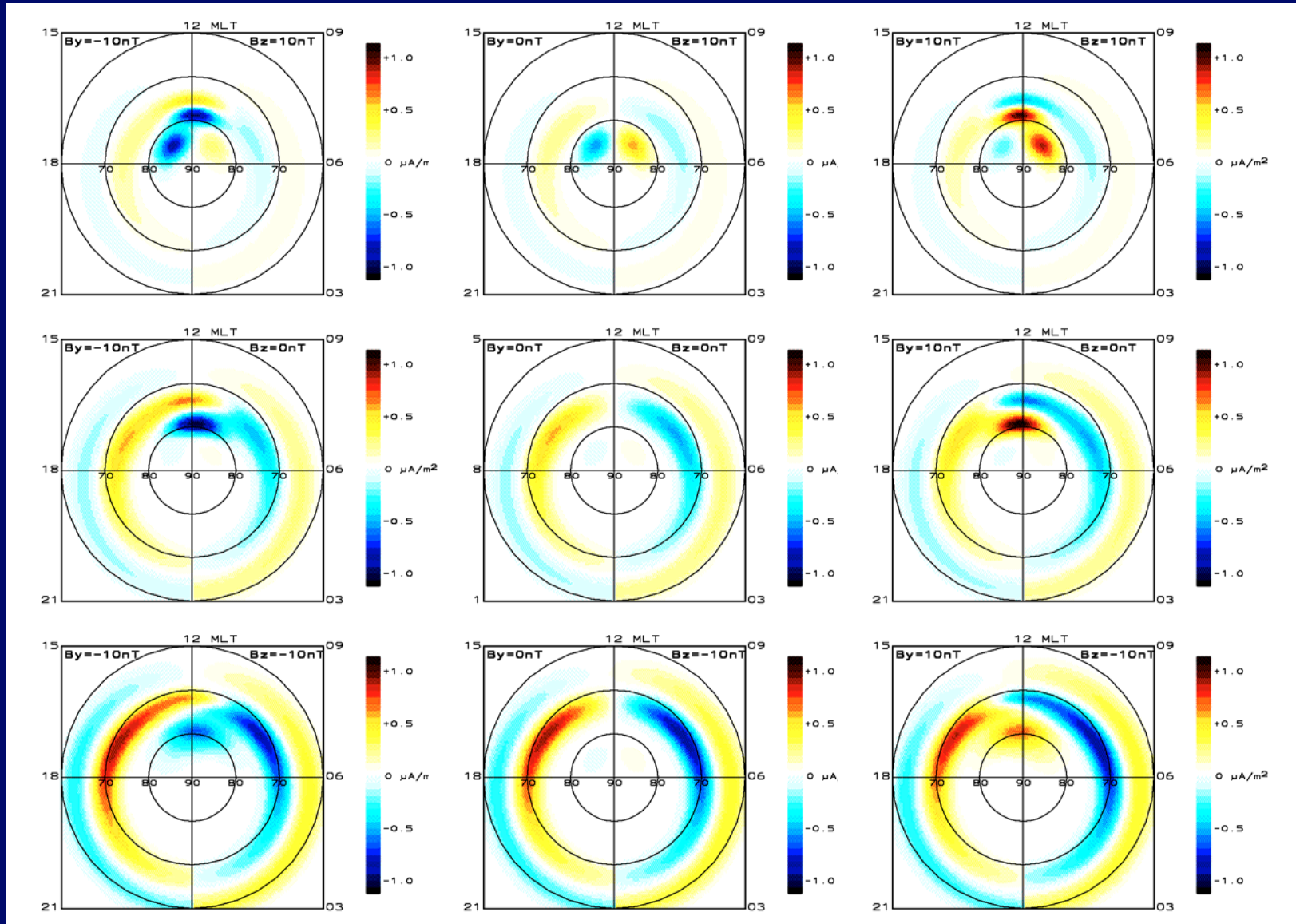
$B_y = 0$ nT

$B_y = +10$ nT

$B_z = +10$ nT

$B_z = 0$ nT

$B_z = -10$ nT



How to determine FAC from satellite magnetometer measurements?

- (a) Inference from a single satellite pass
- (b) Spatial interpolation of data accumulated over many orbits
- (c) Other methods for smoothing 2-D data collections
 - Spherical Cap Harmonic Analysis
 - Spherical harmonic expansion of certain suitable potentials (e.g., magnetic Euler potential)

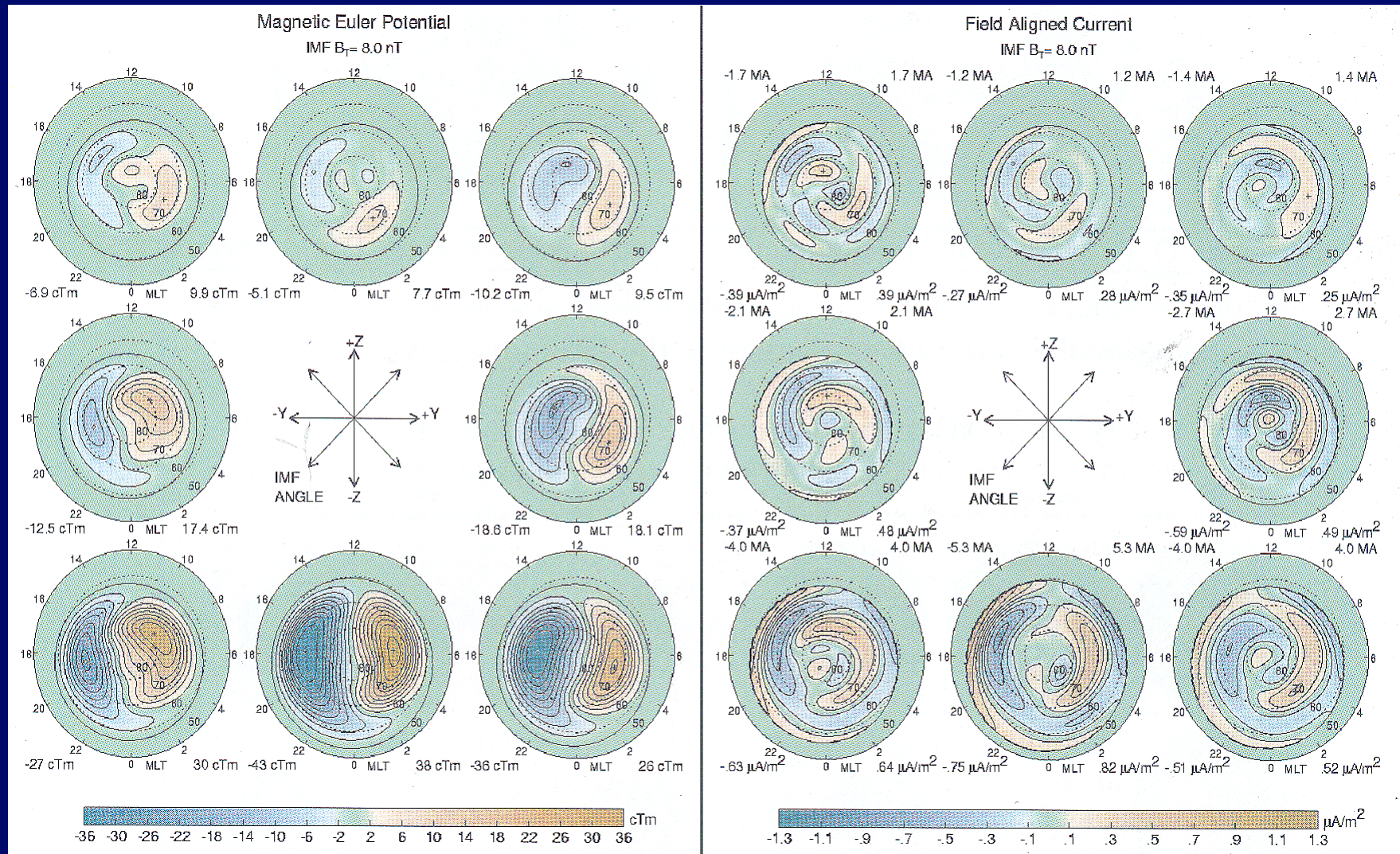
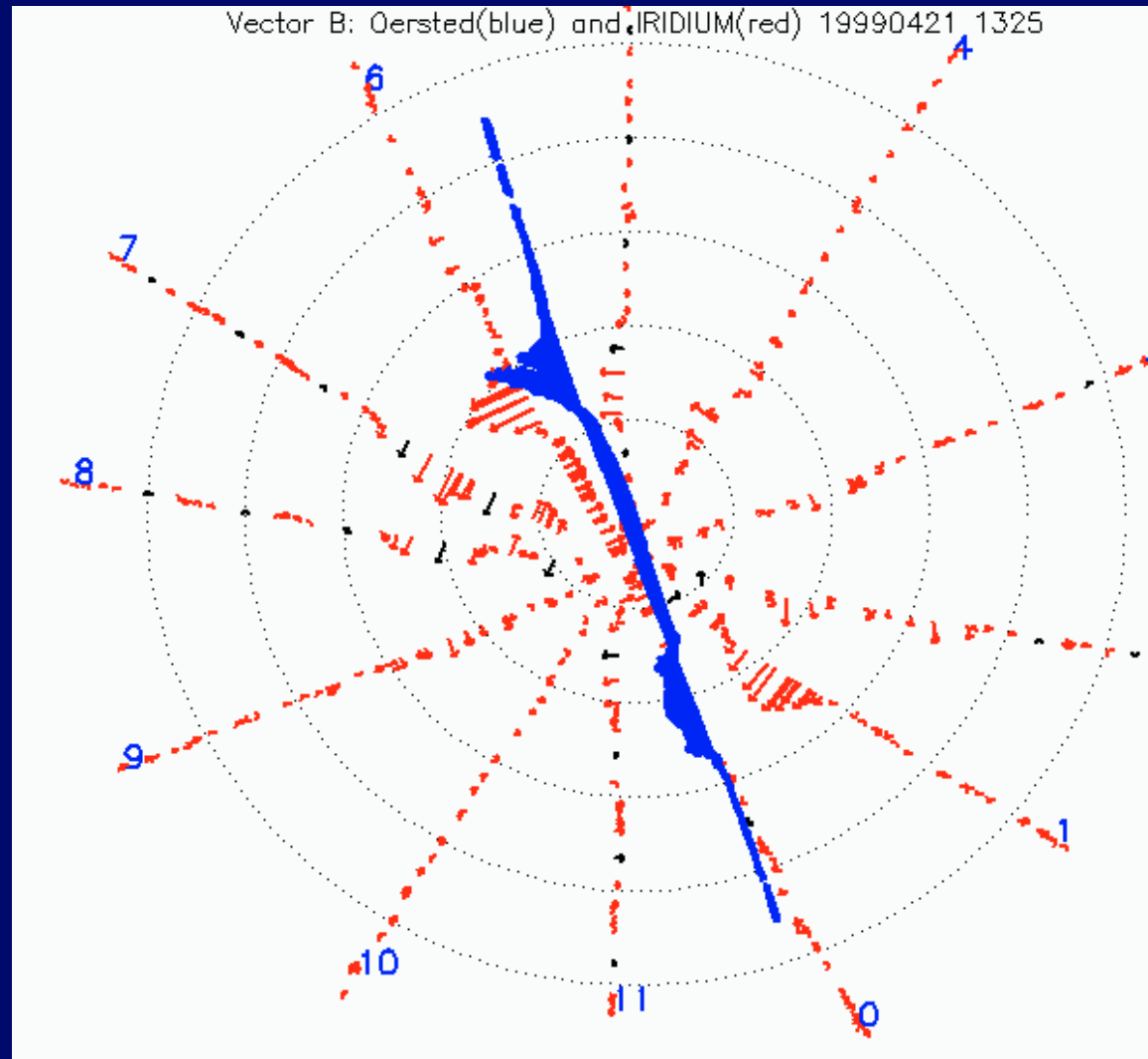


Figure from Weimer, OIST-4 Proceedings (Copenhagen 2003)

How to determine FAC from satellite magnetometer measurements?

- (a) Inference from a single satellite pass
- (b) Spatial interpolation of data accumulated over many orbits
- (c) Other methods for smoothing 2-D data collections
 - Spherical Cap Harmonic Analysis
 - Spherical harmonic expansion of certain suitable potentials (e.g., magnetic Euler potential)
- (d) Taking advantage of a multi-satellite constellation

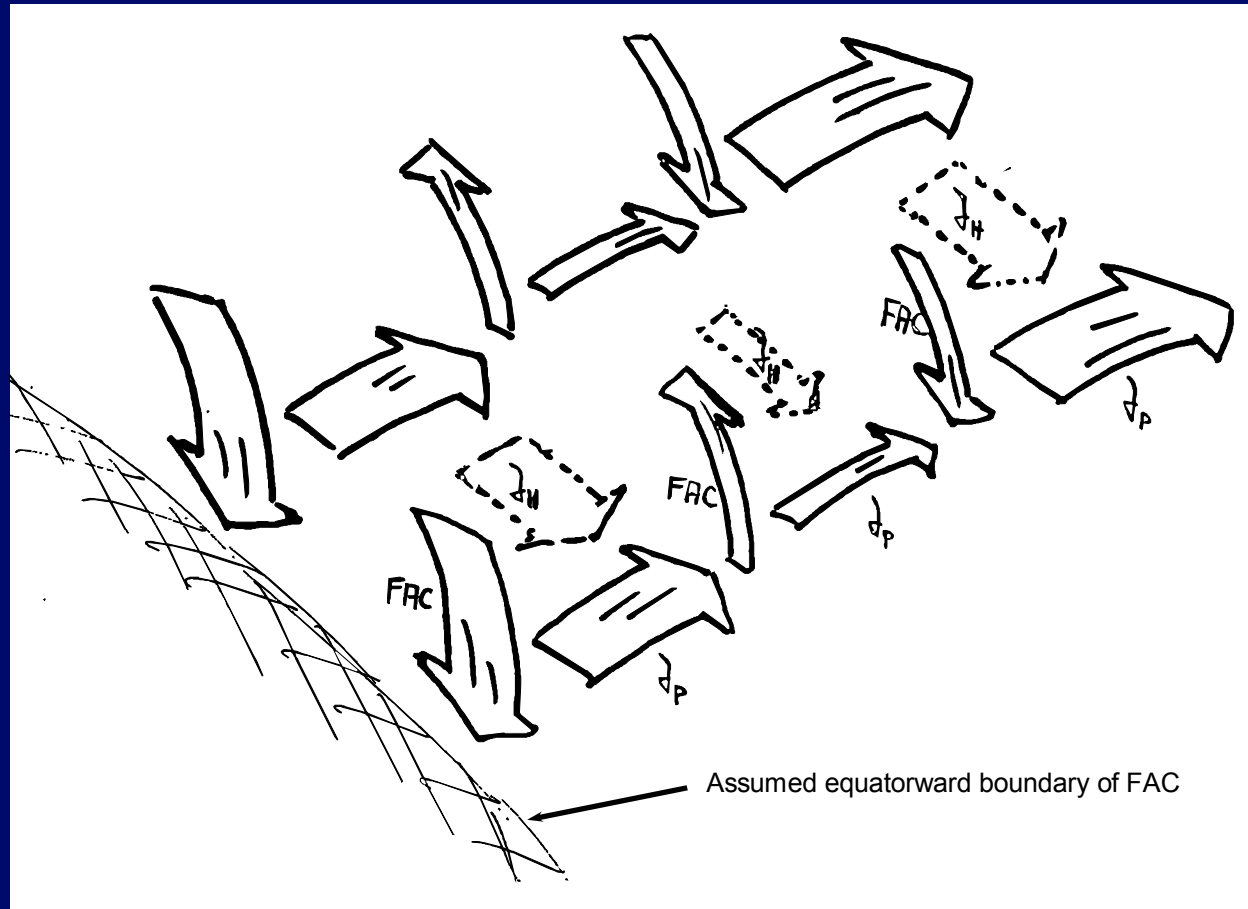
Coincident magnetic field observations from 6 Iridium and the Ørsted satellite



Magnetospheric current system

(magnetosphere-ionosphere coupling currents)

- (1) The large-scale current system
- (2) Field-aligned currents (FAC)
- (3) Ionospheric currents and convection



Sequence of uniform, infinitely extended high-latitude sheet currents

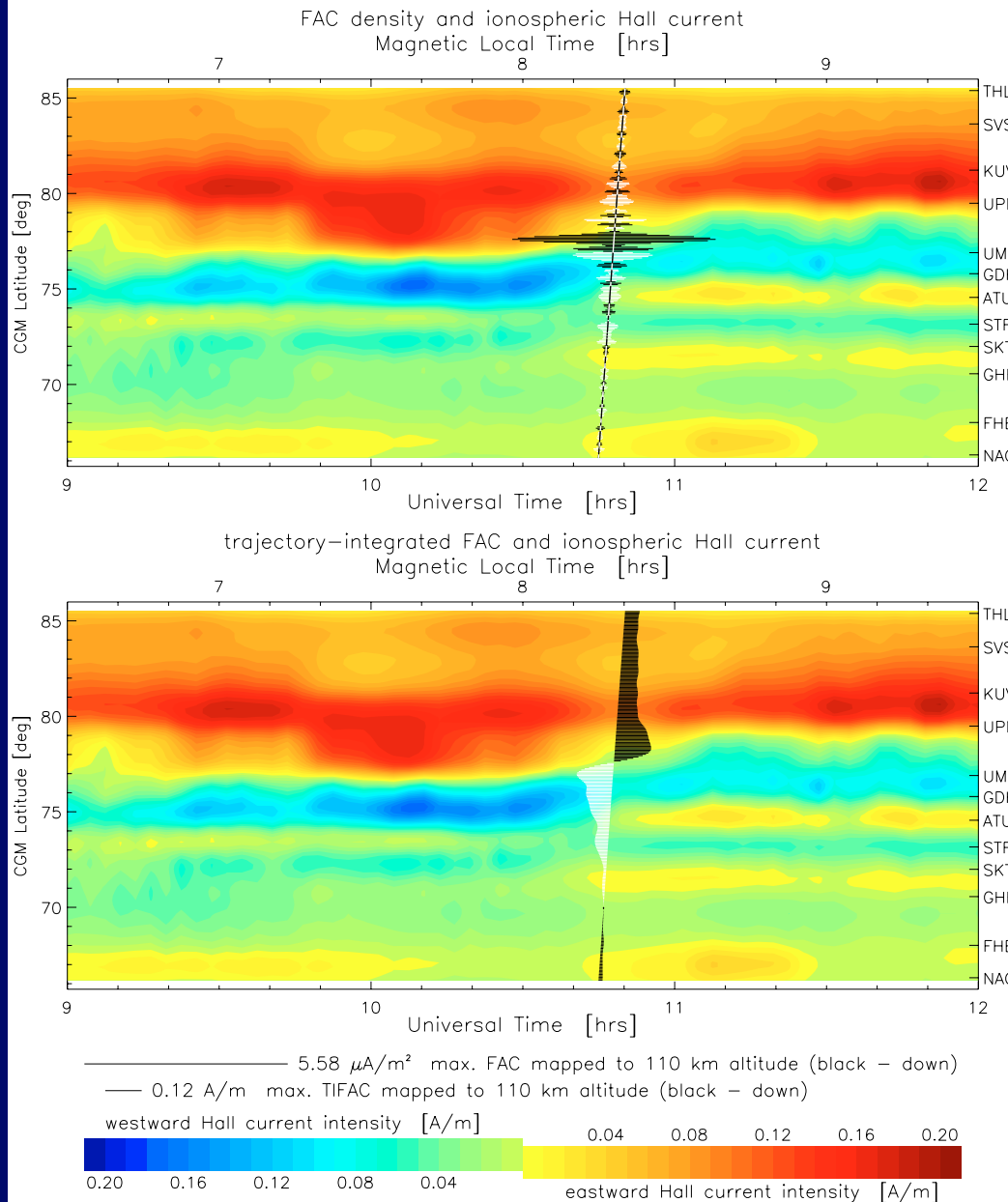
field-aligned currents (FAC)

ionospheric Pedersen currents (J_P)

ionospheric Hall currents (J_H)

Trajectory-integrated FAC (TIFAC) \approx Pedersen current

OERSTED AND GREENLAND WEST COAST MAGNETOMETERS, 2000-04-14



Synoptic view of FAC and ionospheric equivalent current

Ionospheric equivalent current across the magnetometer chain

VS

CGML and time (color-coded)

Ørsted trajectory with FAC density (top panel)

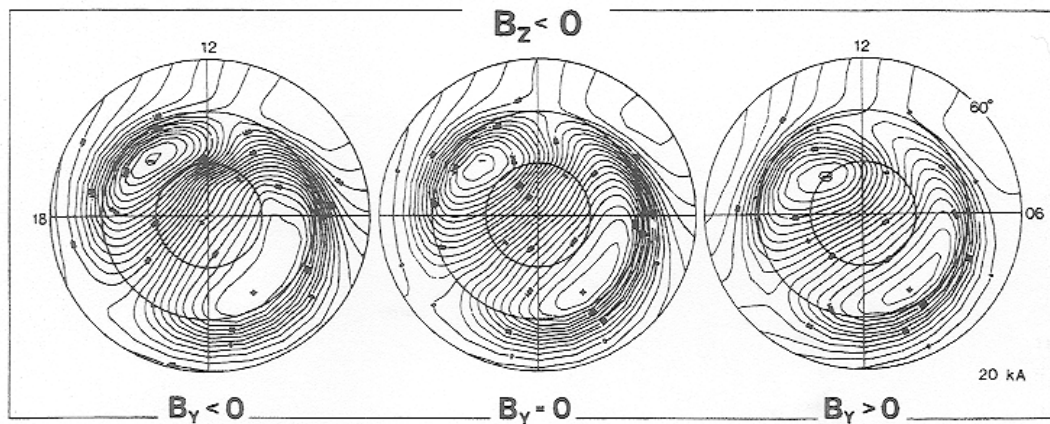
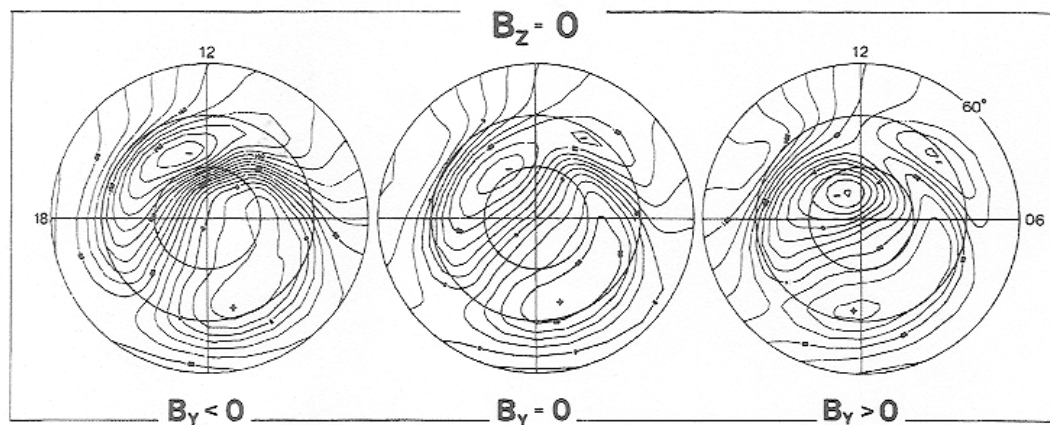
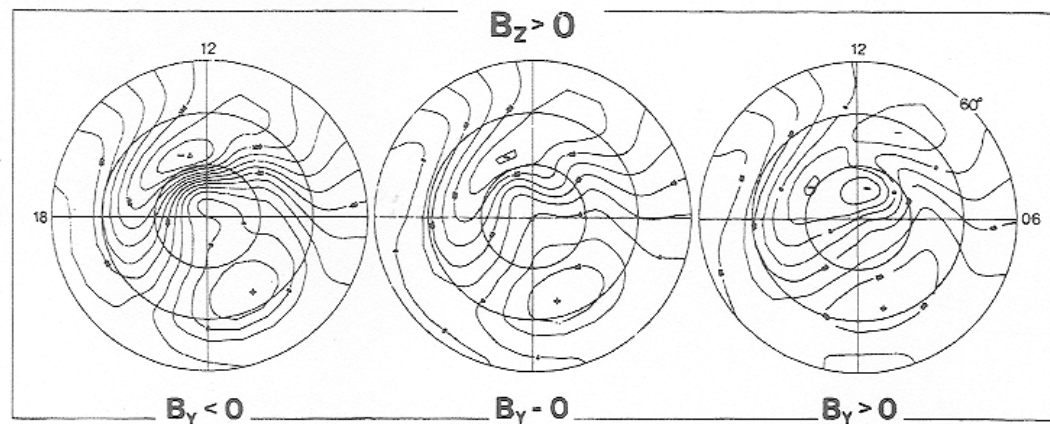
and trajectory integrated FAC (bottom panel)

white -- upward

black – downward

Input	Assumptions	Output	Name of method	Remarks
\vec{B}_G	-	$\vec{J}_{eq, Ion}, \vec{J}_{eq, Int}$	field continuation, field separation	forward method no true currents, no FACs
\vec{B}_G	Σ_H, Σ_P	$\vec{E} (\Phi_E), \vec{J}, j_{\parallel}$	KRM	forward method boundary conditions critical if non-global
$\vec{B}_G,$ $\{\vec{E} (\Phi_E),$ satellite data}	Σ_H, Σ_P	$\{\vec{E} (\Phi_E)\}, \vec{J}, j_{\parallel}$	AMIE, “three- dimensional modeling”	optimisation method (least square or “trial and error”); data need not to cover whole analysis area
\vec{B}_G, \vec{E} ----- j_{\parallel}, \vec{E}	$\alpha = \Sigma_H / \Sigma_P$	$\Sigma_H, \Sigma_P, \vec{J}, j_{\parallel}$	method of characteristics (JEQ-based) ----- (FAC-based)	forward method α assessible from ASC or \vec{B}_G
j_{\parallel}	Σ_H, Σ_P	$\vec{E} (\Phi_E), \vec{J}$	-	forward method boundary conditions critical if non-global
$\vec{B}_G, j_{\parallel}, \{\vec{E}\}$	-	$\vec{J}, \{\Sigma_H, \Sigma_P\}$	elementary current method	forward method

Table from Amm et al., OIST-4 Proceedings (Copenhagen 2003)



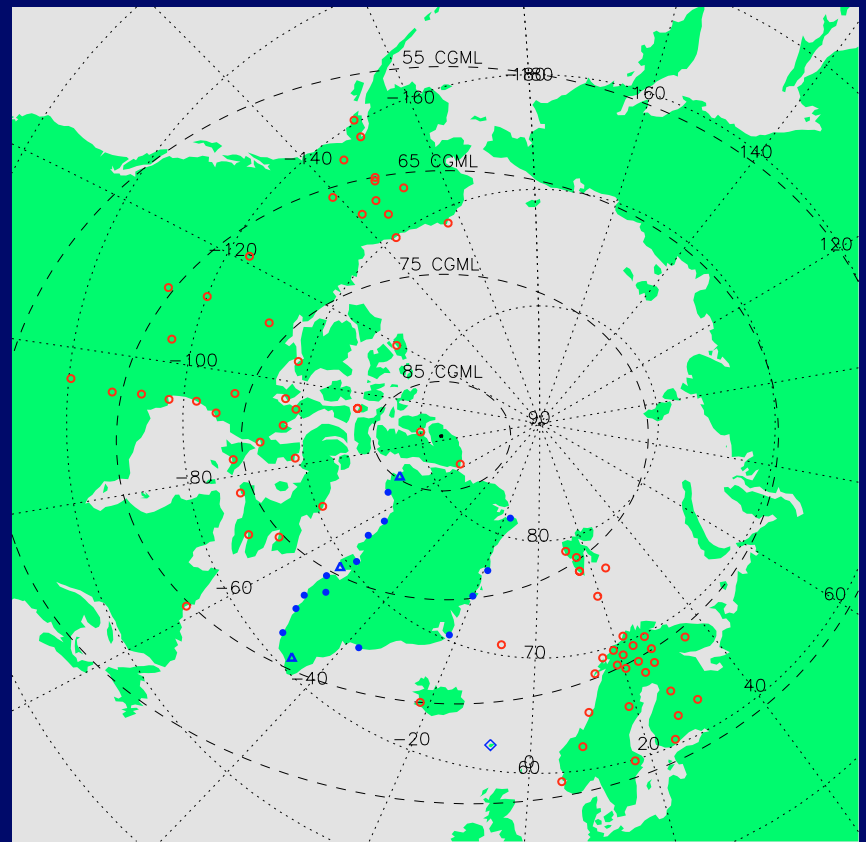
Equivalent ionospheric currents for various combinations of the interplanetary magnetic field components B_z and B_y

Figure from Friis-Christensen (1985)



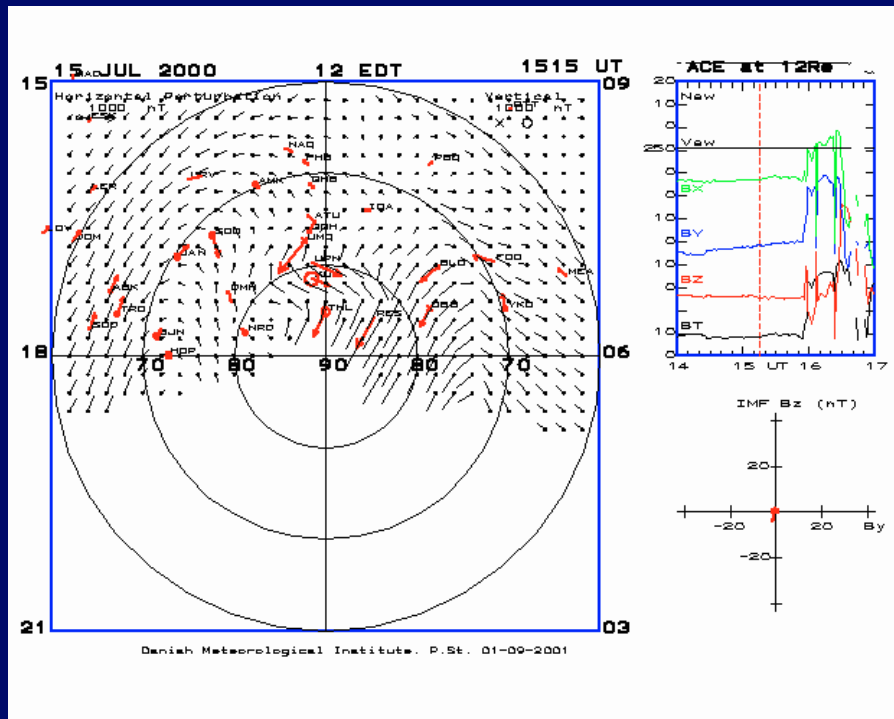
HF coherent scatter radar coverage of the northern hemispheric ionosphere

figure by JHU / APL



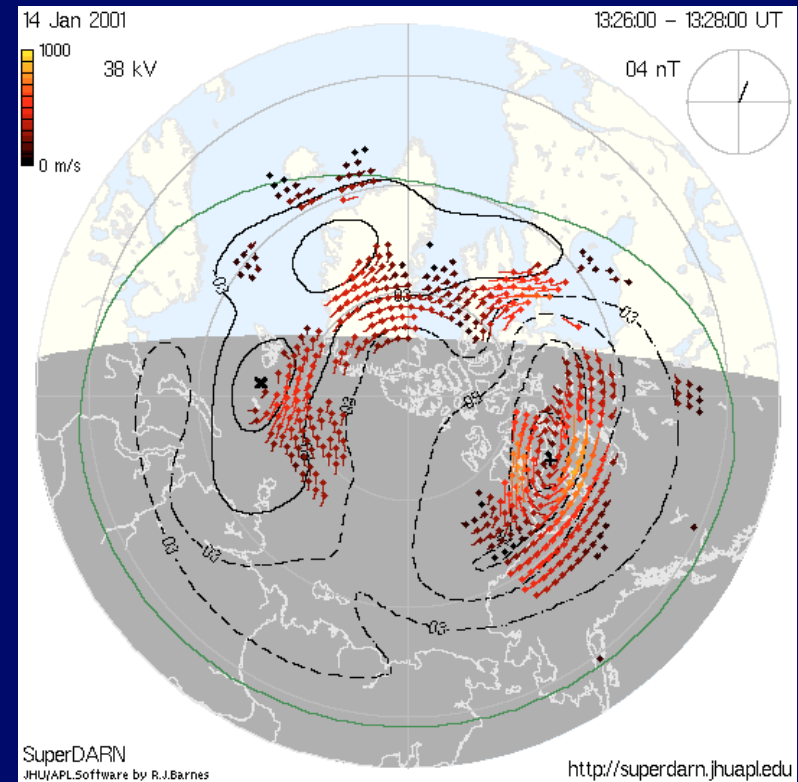
Magnetometer sites

blue – DMI stations
red – other institutions



Ionospheric convection inferred from northern hemisphere ground-based magnetometers

Frame from a convection movie, Peter Stauning, DMI



SuperDARN sample convection pattern

Figure from Opgenoorth et al., 2001

Geomagnetic field models and indices



Jurgen Watermann

Geomagnetism and Space Physics Program

Danish Meteorological Institute

May 2006



*The support provided by COST, COST Action 724
and the Abdus Salam ICTP is gratefully acknowledged*

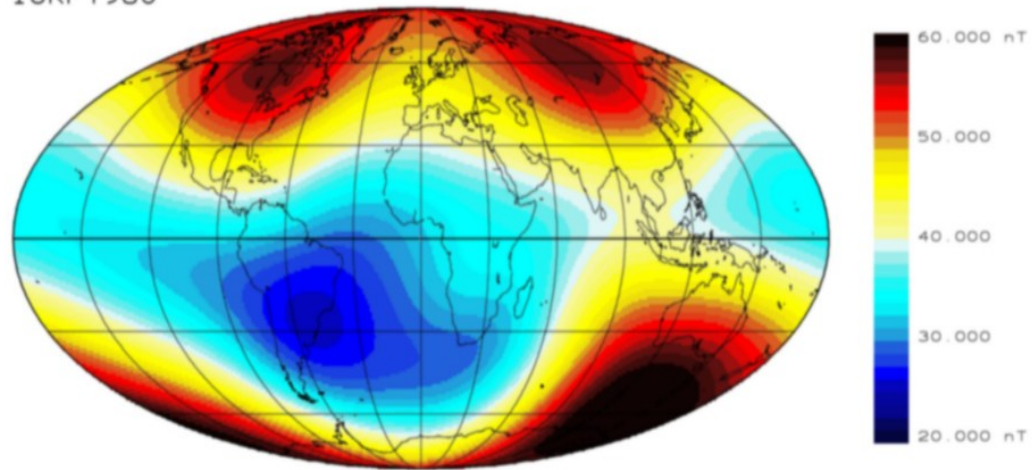
Geomagnetic and magnetospheric field models

- (1) IGRF geomagnetic model

IGRF – a model of the geomagnetic main field

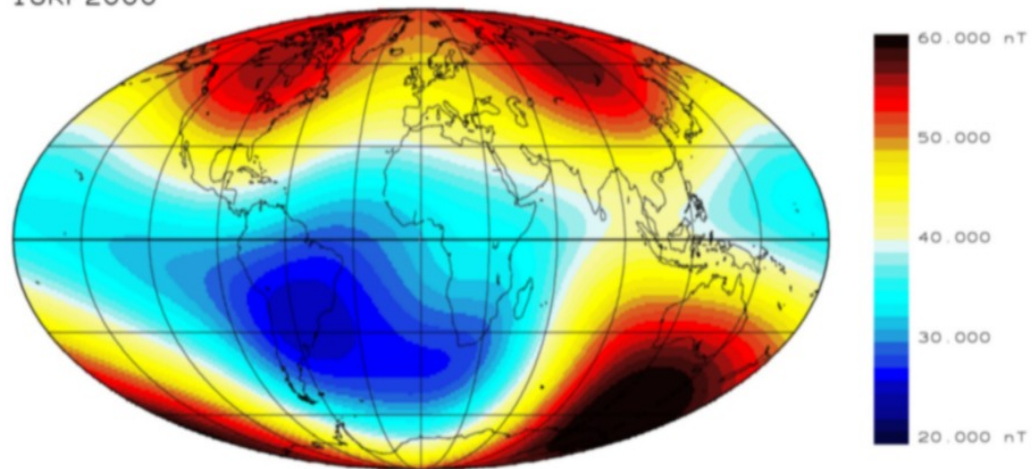
- only quiet days are considered
- external contributions are removed to the most possible extent
- sources are considered solely internal
 - ⇒ magnetic field at and above the Earth's surface is a potential field
- the magnetic field potential is expanded in spherical harmonics
 - up to year 2000 with order and degree 10
 - after year 2000 with order and degree 13
- the coefficients of the expansion are updated every 5 years in order to account for the secular variation

IGRF 1980



IGRF 1980 based
on MAGSAT data
total field strength

IGRF 2000



IGRF 2000 based
on Ørsted data
total field strength

Geomagnetic and magnetospheric field models

- (1) IGRF geomagnetic model
- (2) Tsyganenko magnetospheric model

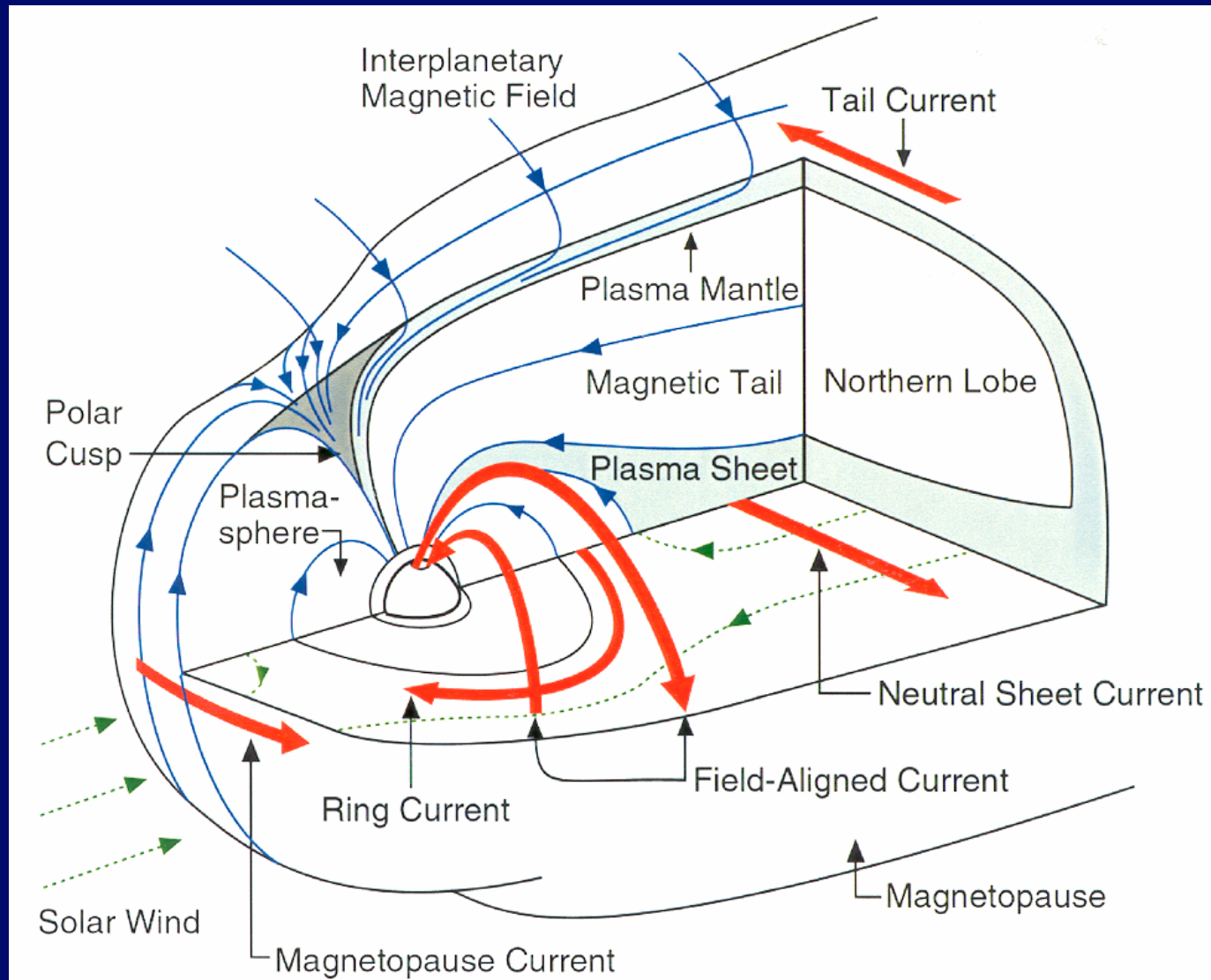
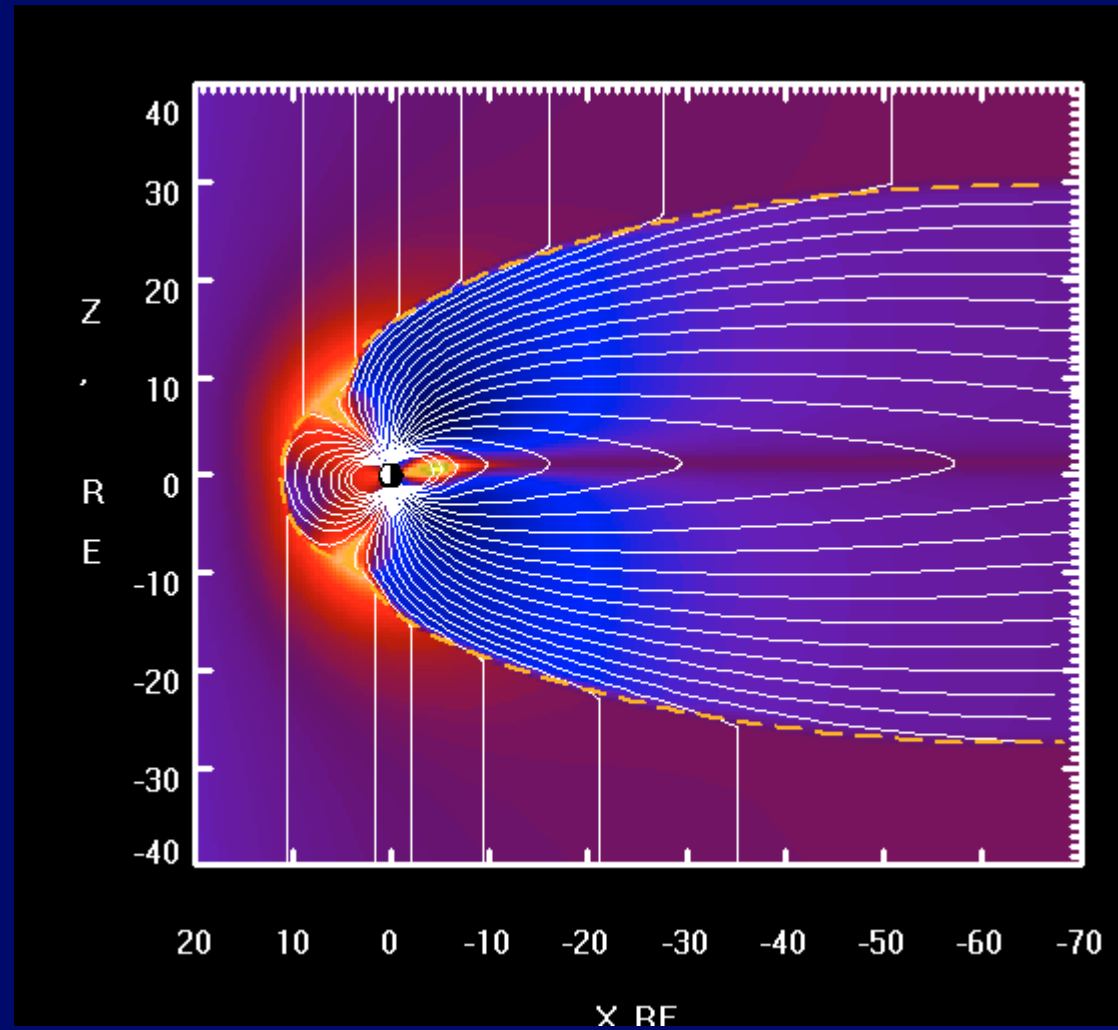


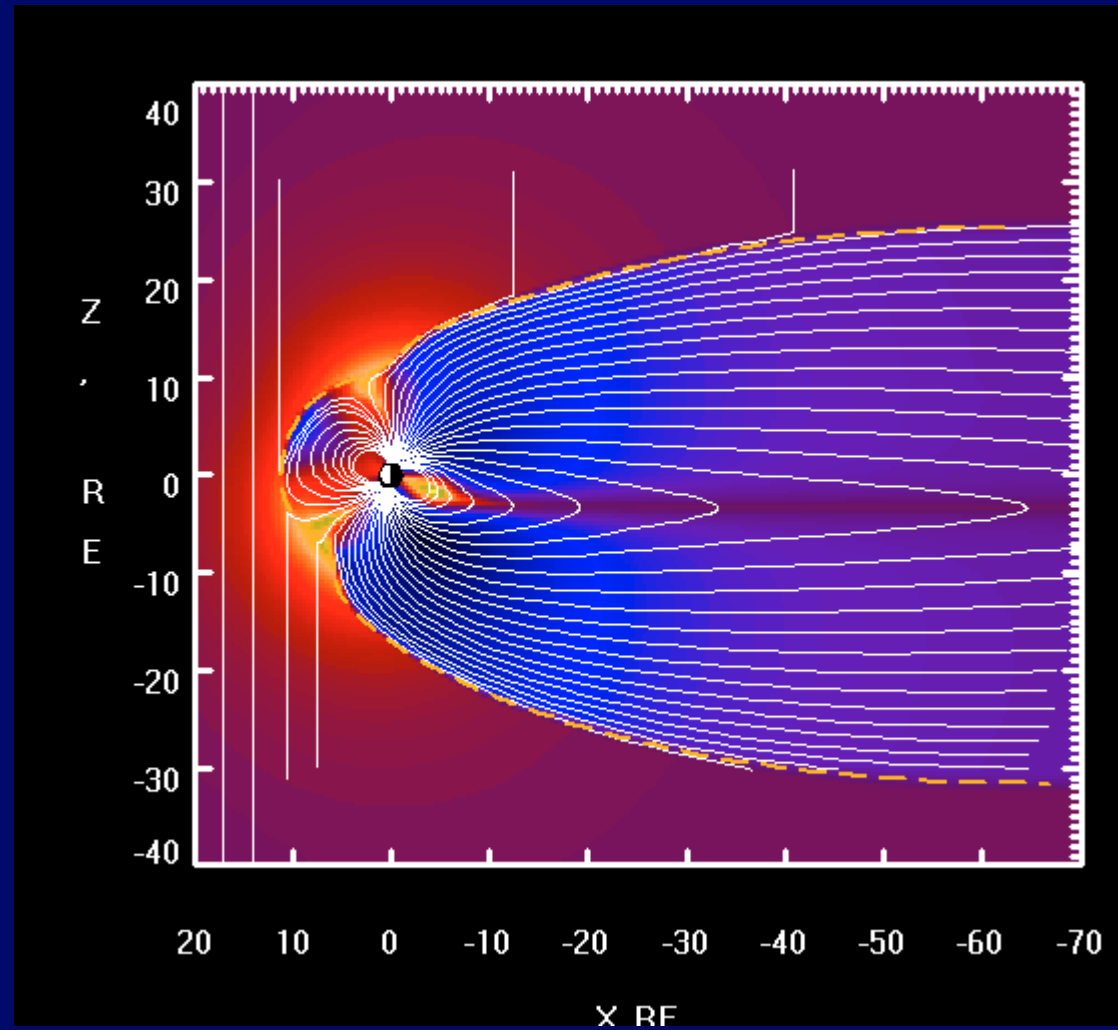
Figure from "Report from the first Geospace Environment Modeling (GEM) Campaigns: 1991-1997"

Tsyganenko 95 model



Animation by N. Tsyganenko

Tsyganenko 95 model



Animation by N. Tsyganenko

IAGA approved magnetic indices

for monitoring magnetospheric and ionospheric electric currents

description available at the ISGI home page

<http://www.cetp.ipsl.fr/~isgi/homepage1.htm>

Dst

quantitative magnetic storm index (physical unit nT)

« Storm is an interval of time when a sufficiently intense and long-lasting interplanetary convection electric field leads, through a substantial energization in the magnetosphere-ionosphere system, to an intensified ring current strong enough to exceed some key threshold of the quantifying storm time index Dst. »

Gonzalez et al., 1994

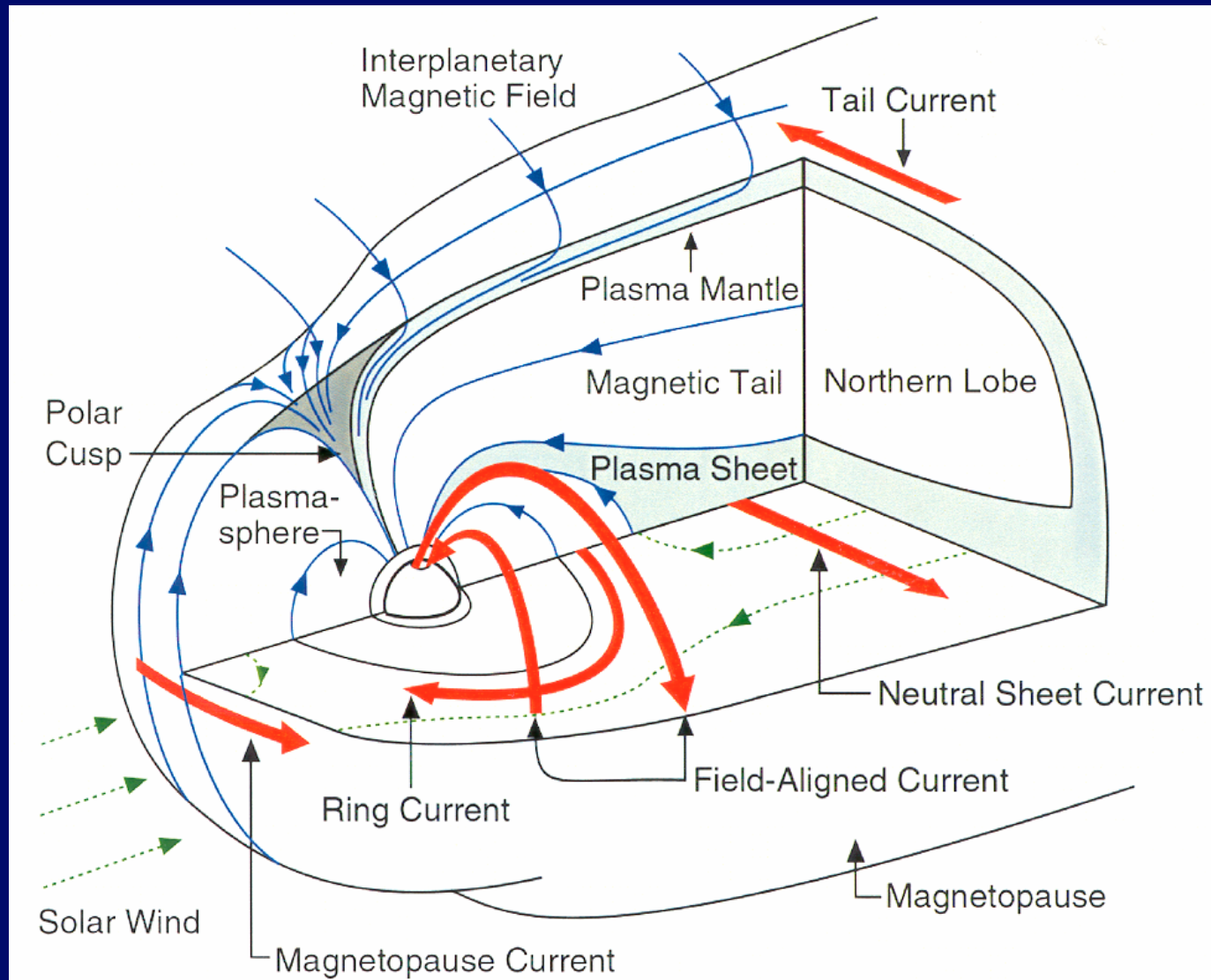


Figure from "Report from the first Geospace Environment Modeling (GEM) Campaigns: 1991-1997"

Details of the Dst index

Objective:

quantify the axisymmetric magnetic field of magnetospheric currents (primarily the ring current, secondarily magnetopause currents)

Contributors:

four low-latitude (but not equatorial) observatories

Physical quantity:

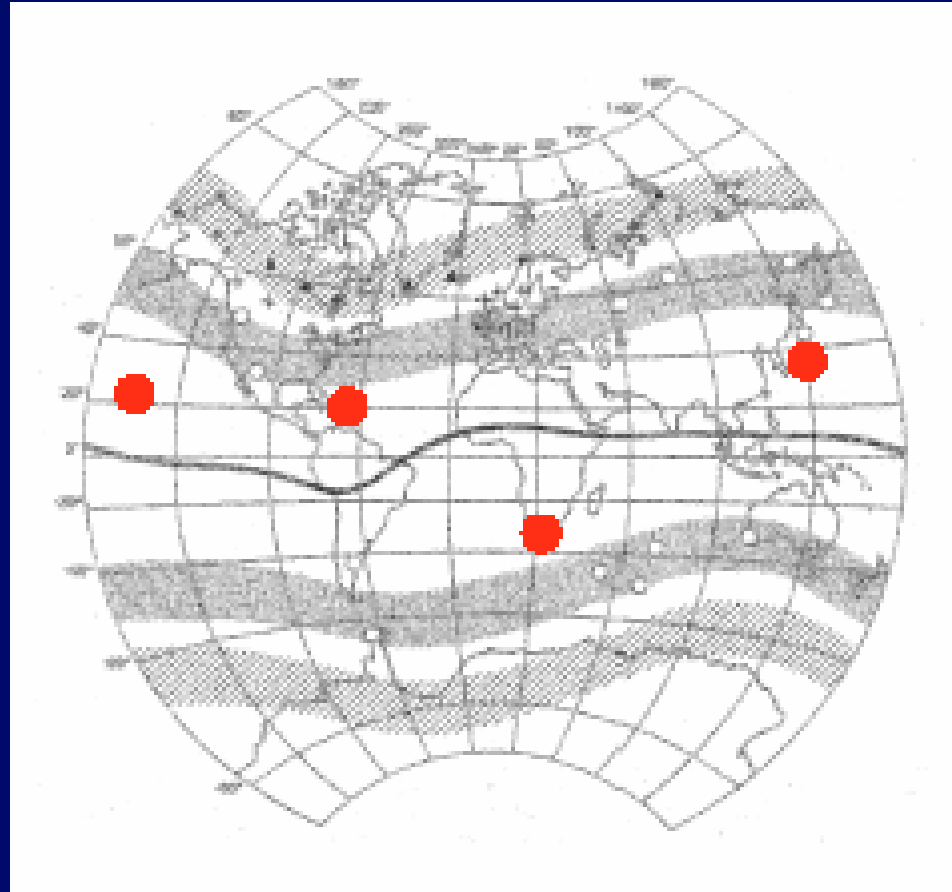
deviation of the horizontal magnetic field component from reference level

Reference level:

quasi-static internal geomagnetic field plus Sq (solar quiet) variation

Method of derivation:

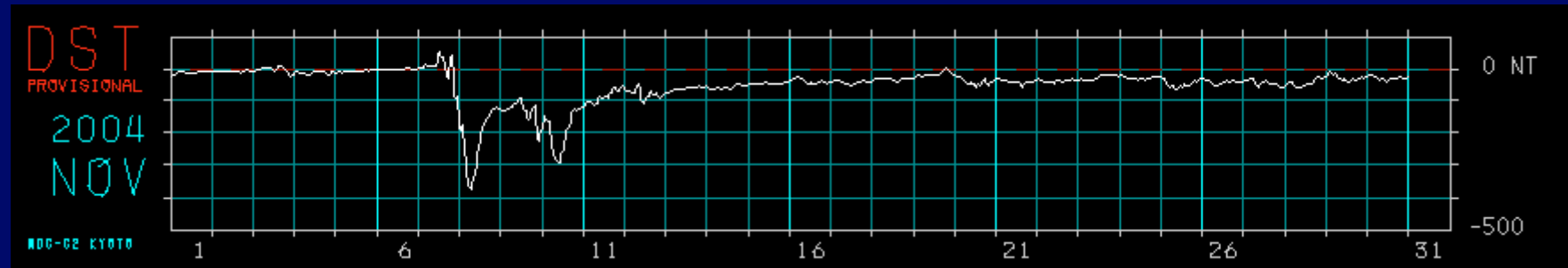
- (1) at each station: $D_k =$ hourly mean of ΔH_k
- (2) $Dst = \{\text{mean of } D_k (k=1, \dots, 4)\} / \{\text{mean of } \cos(\text{dip}_k)\}$



*Figure from Menvielle and Marchaudon (2006)
"Geomagnetic Indices in Solar-Terrestrial Physics and Space Weather"*

Provisional Dst

provided by the World Data Center for Geomagnetism Kyoto (courtesy of T. Kamei)



Storm strength

severe (super storm)

intense (strong)

moderate

Dst exceeds

-200 nT / -250 nT

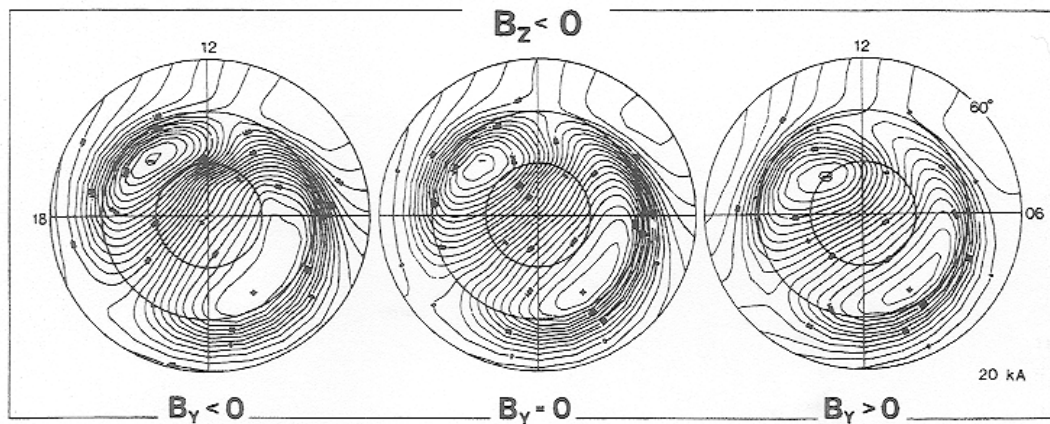
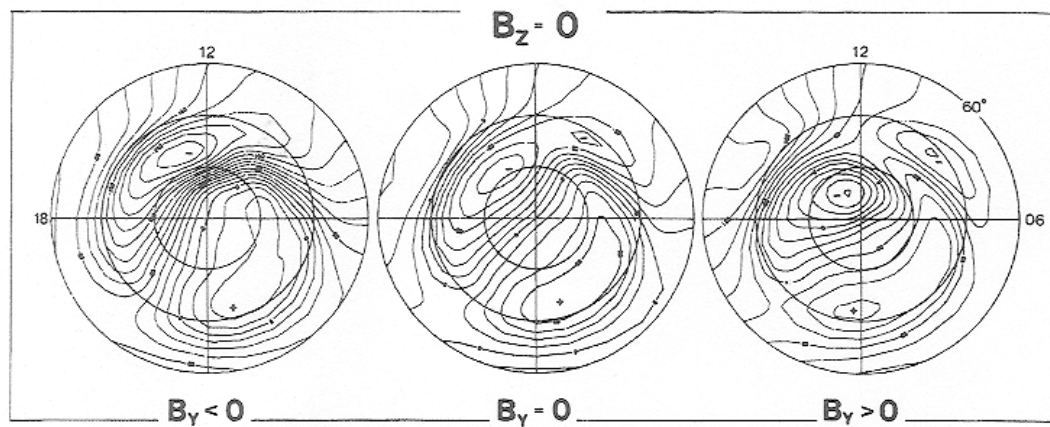
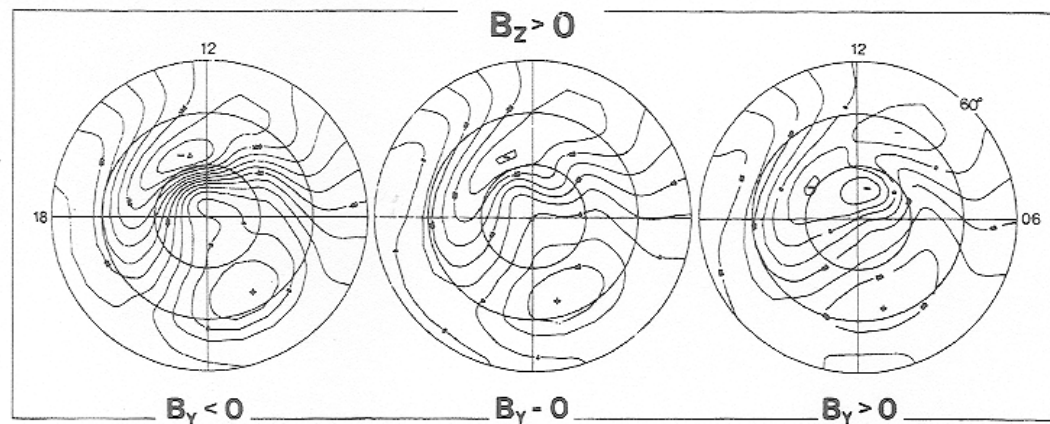
-100 nT

-50 nT

AE (AU, AL, A0)

quantitative auroral activity index (physical unit nT)

The AE group of indices monitors the enhancement of the global auroral ionospheric currents (electrojets) which are mostly associated with magnetosphere-ionosphere coupling through field-aligned currents (R1 and R2 currents).



Equivalent ionospheric currents for various combinations of the interplanetary magnetic field components B_z and B_y

Figure from Friis-Christensen (1985)

Details of the auroral activity indices

Objective:

quantify the magnetic field of the auroral electrojets
(AU signifies the eastward and AL the westward current)

Contributors:

11 observatories in the nominal northern auroral zone

Physical quantity:

deviation of the horizontal magnetic field component from reference level

Reference level:

quasi-static internal geomagnetic field plus Sq (solar quiet) variation

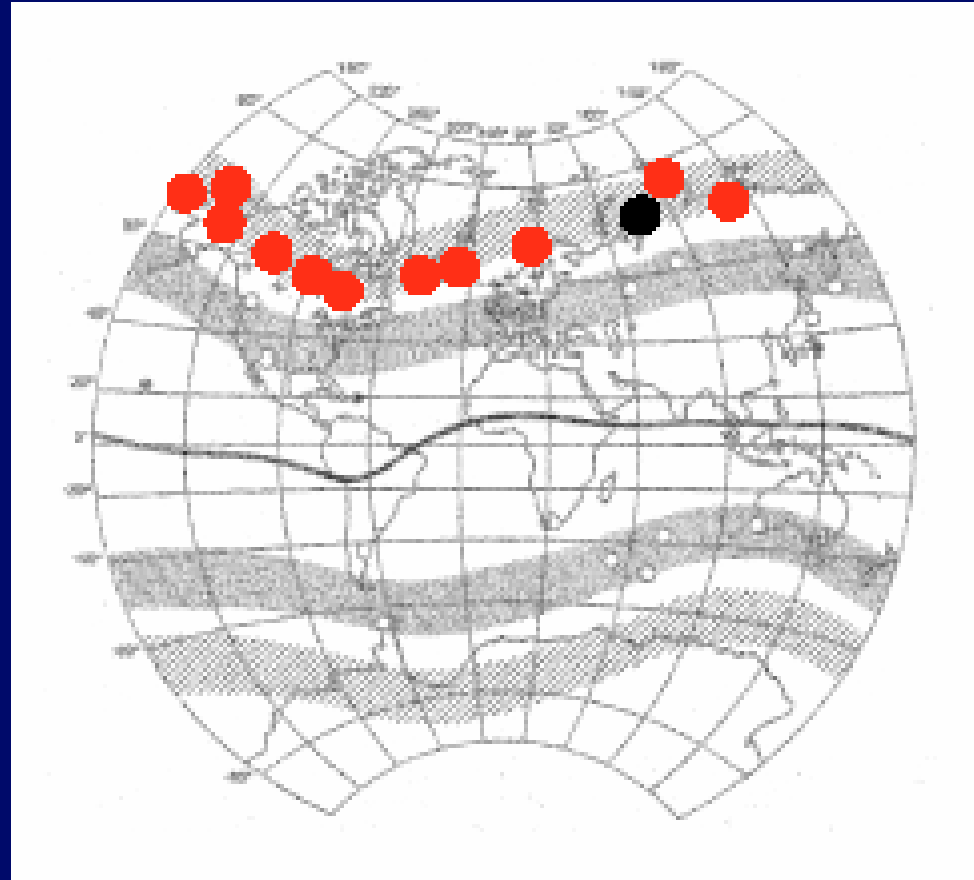
Method of derivation:

AU = largest ΔH recorded at the contributing stations (1-min mean)

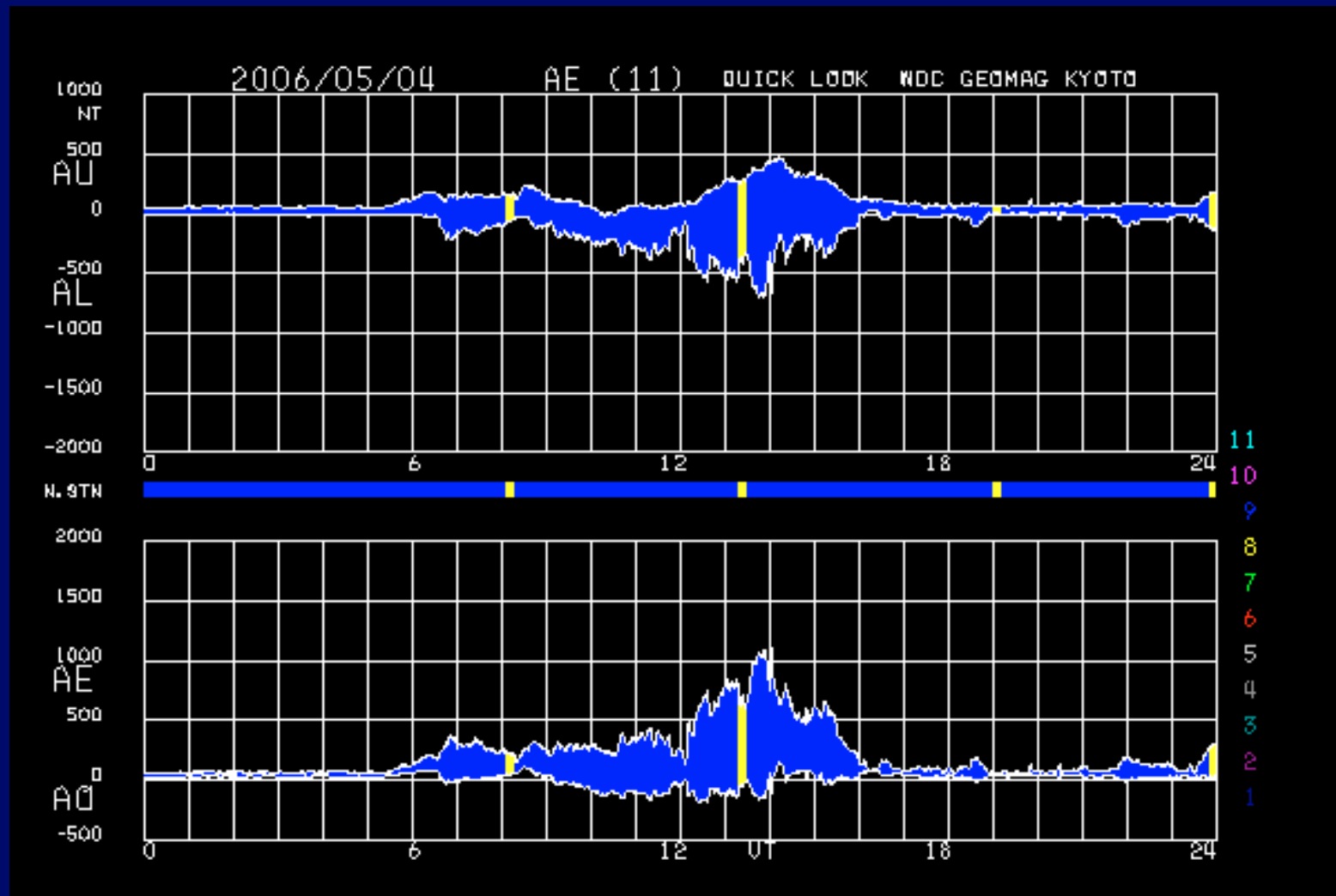
AL = smallest ΔH recorded at the contributing stations (1-min mean)

AE = AU – AL

A0 = (Au + AL) / 2



*Figure from Menvielle and Marchaudon (2006)
"Geomagnetic Indices in Solar-Terrestrial Physics and Space Weather"*



Real-time (quicklook) Auroral Activity indices

provided by the World Data Center for Geomagnetism Kyoto (courtesy of T. Kamei)

Potential problem (in fact, real problem)

When the cross polar cap electric potential is high the auroral oval expands and the centers of the electrojets move equatorward. The AE stations, per definition located in the nominal auroral oval, may in fact be found poleward of the actual auroral oval.

Kp and related indices

semi-quantitative global general activity indices

(K^* dimensionless, a^* in nT scale)

The Kp group of indices monitors to some extent the overall magnetospheric activity (energy content of the magnetosphere)

Details of the Kp and related indices

Objective:

quantify the overall geomagnetic activity

Contributors:

13 observatories at subauroral latitudes (Kp and ap)

23 observatories at subauroral latitudes (Km and related)

Physical quantity:

range value of the horizontal magnetic field component

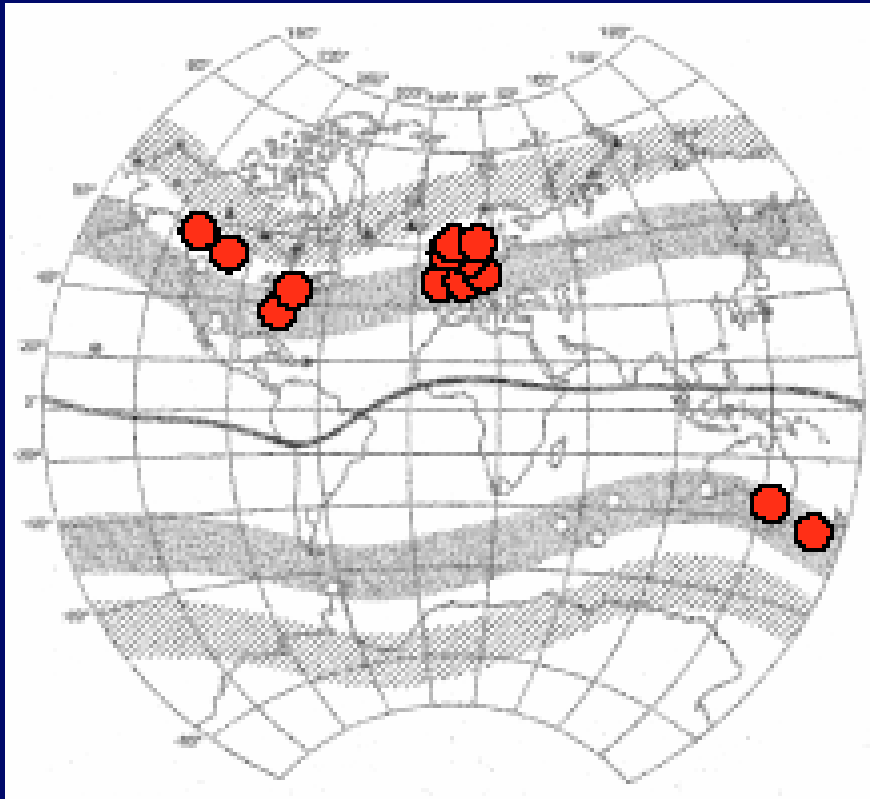
Reference level:

not required but Sq (solar quiet) variation is subtracted

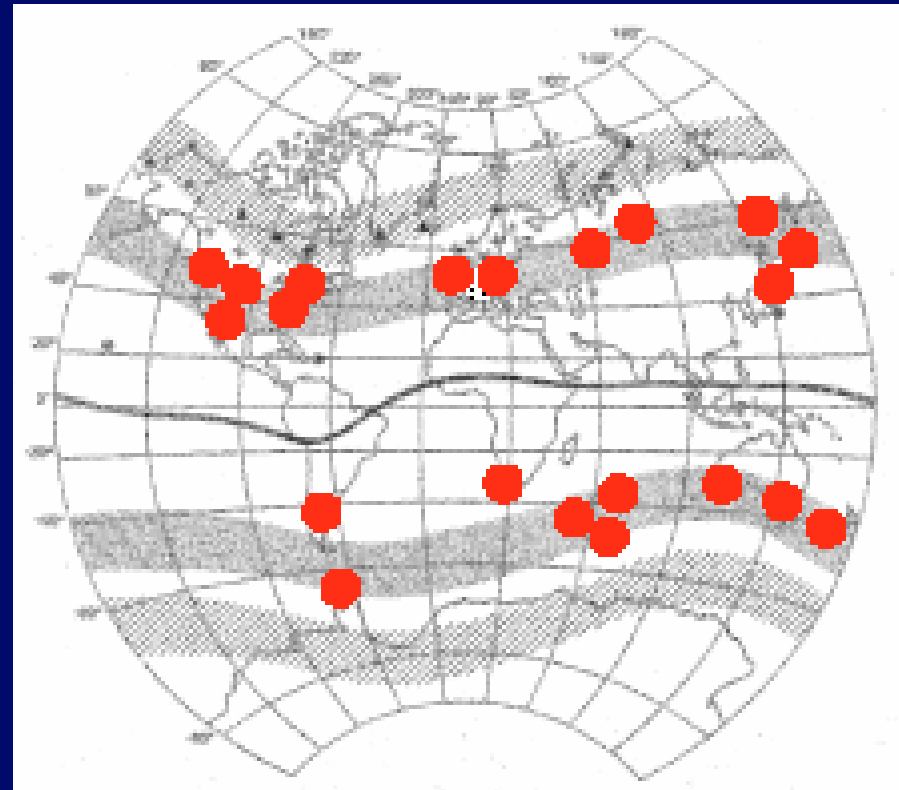
Method of derivation:

3-hour range (max–min) of the horizontal magnetic field
a bit awkward since it involves empirical scaling

Kp (ap) observatories

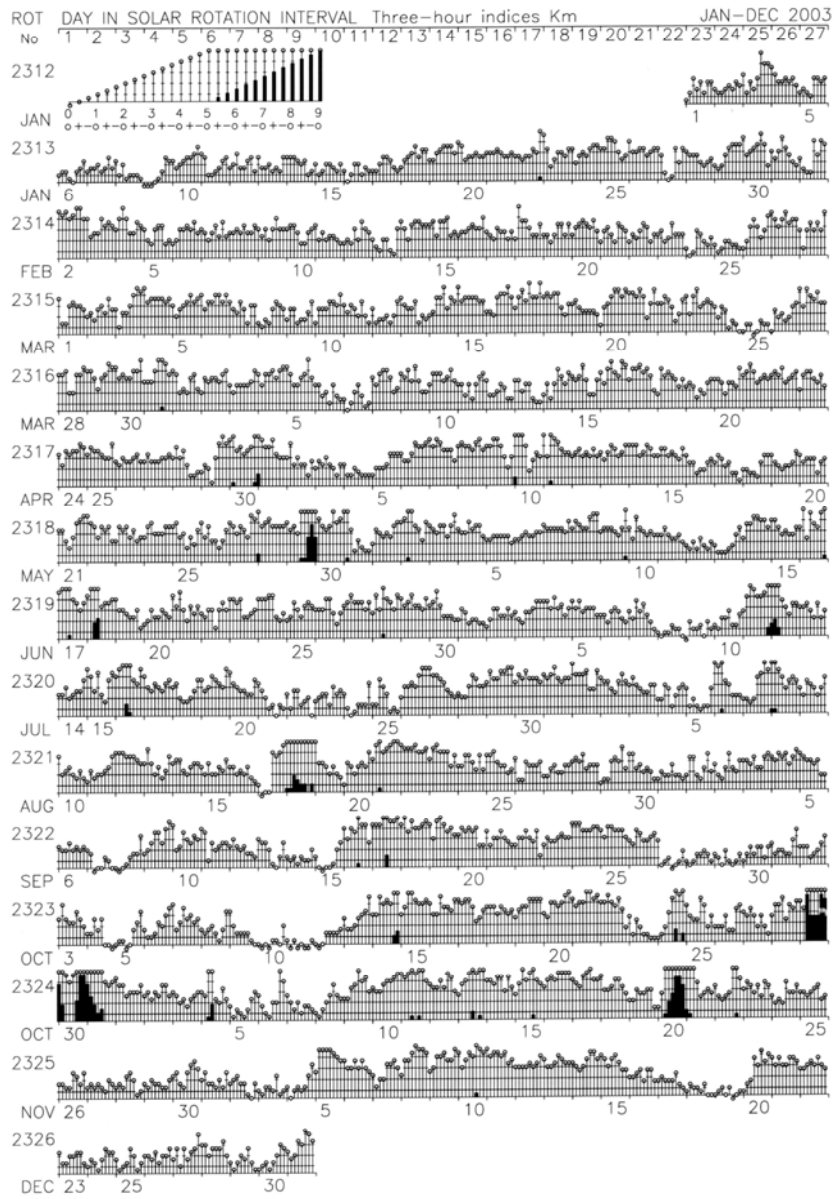


Km (an, as, am) observatories



*Figures from Menvielle and Marchaudon (2006)
"Geomagnetic Indices in Solar-Terrestrial Physics and Space Weather"*

MUSICAL DIAGRAM OF Km 2003



Km "musical diagram"
provided by the GFZ Potsdam

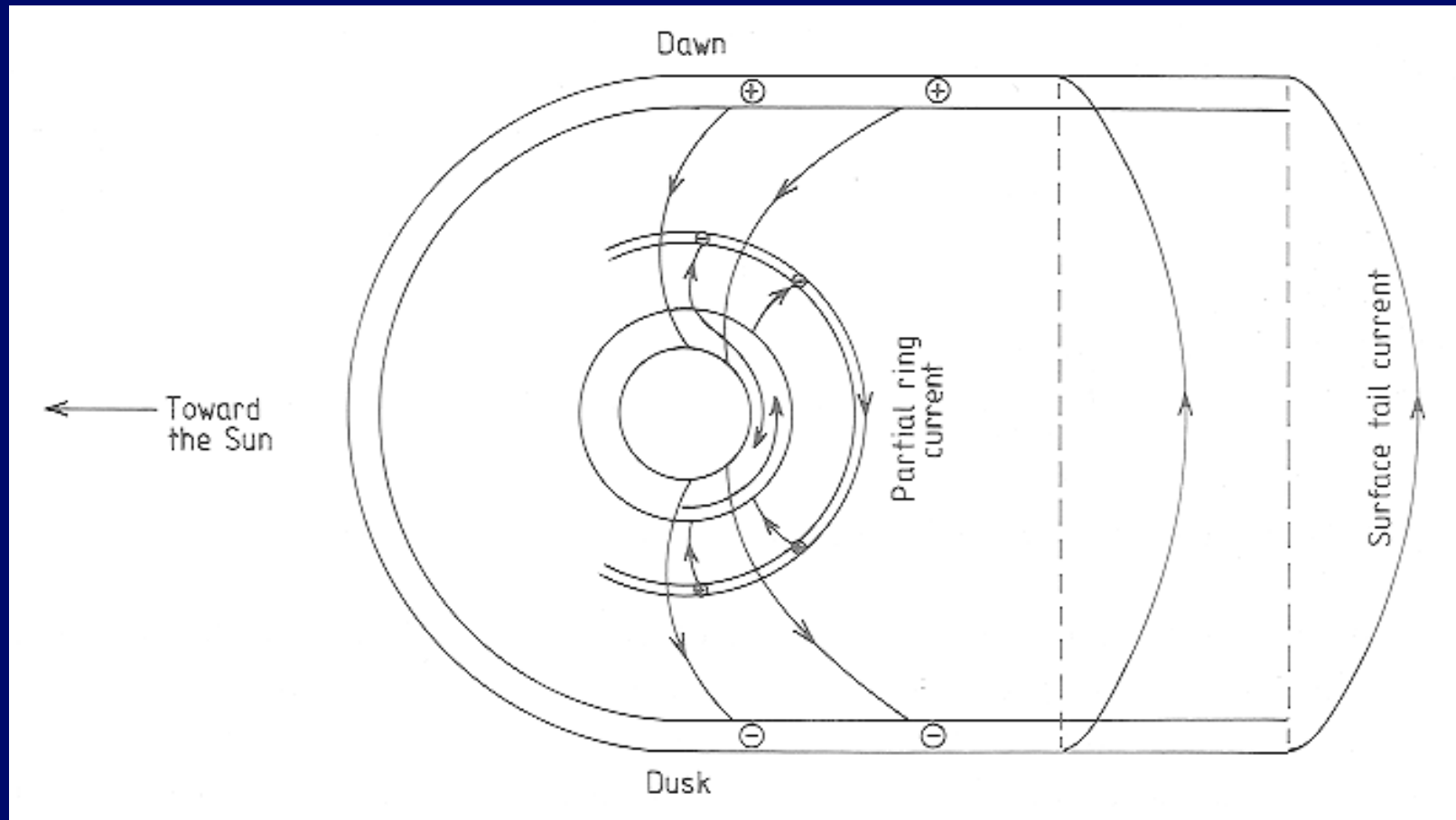
The Kp group is not well suited for space weather applications

- the indices represent a three-hour interval
- they are range values rather than absolute values
- they possess a maximum class which cannot properly account for extreme events
- the Kp (and ap) indices stem from a very unevenly distributed sites

Magnetic indices not (yet) IAGA approved

SYM and ASY

quantitative Dst related indices



Closure of near-Earth magnetospheric current system

Figure from Brekke "Physics of the Upper Atmosphere" (1997)

Details of the SYM-H and ASY-H indices

Objective:

improve upon the Dst index through

- increasing the number of stations
- monitoring the symmetric and asymmetric part of the ring current
- increasing the time resolution from 1 hour to 1 minute

Contributors:

six low- and mid-latitude observatories resp. observatory pairs

Physical quantity:

deviation of the horizontal magnetic field component from reference level

Reference level:

quasi-static internal geomagnetic field plus Sq (solar quiet) variation

Method of derivation:

as for Dst computation but based on 1-min averages

Characteristics:

SYM-H is essentially Dst on a 1-min basis

ASY-H is sensitive to substorms (like AE)

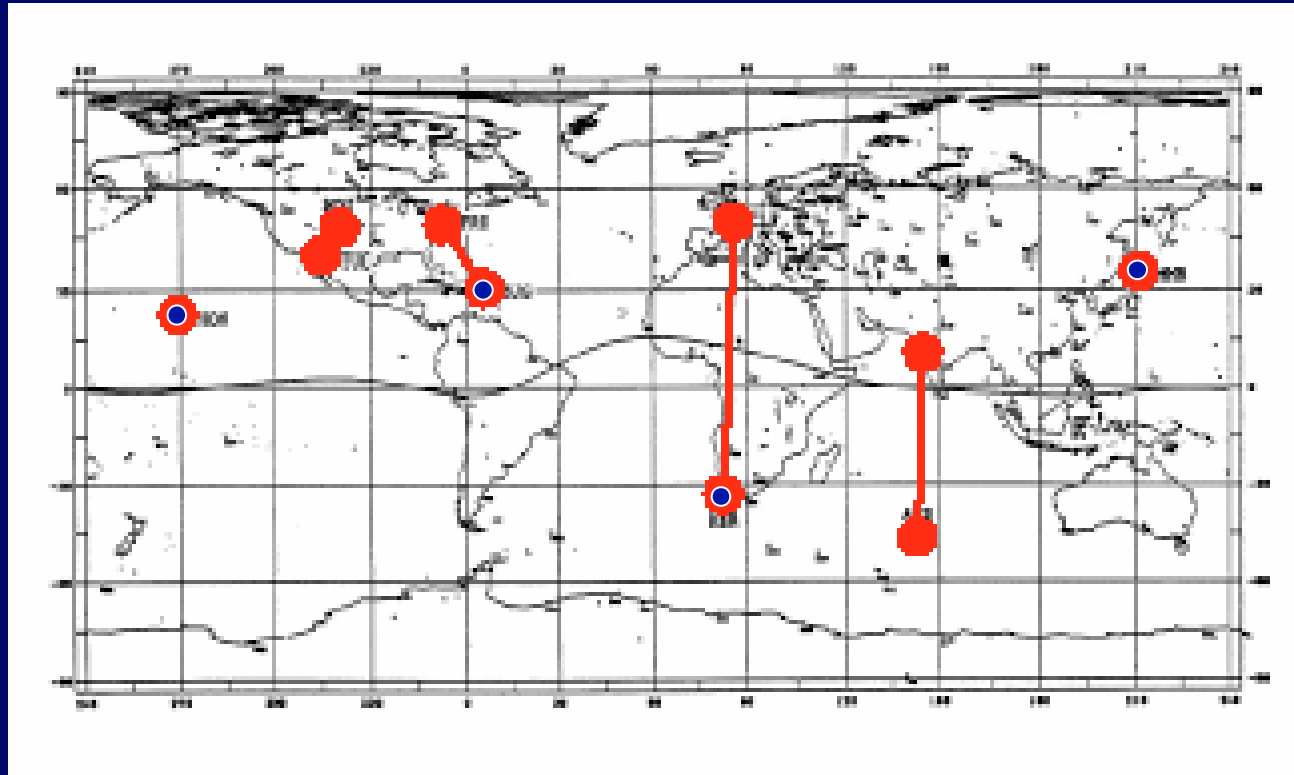
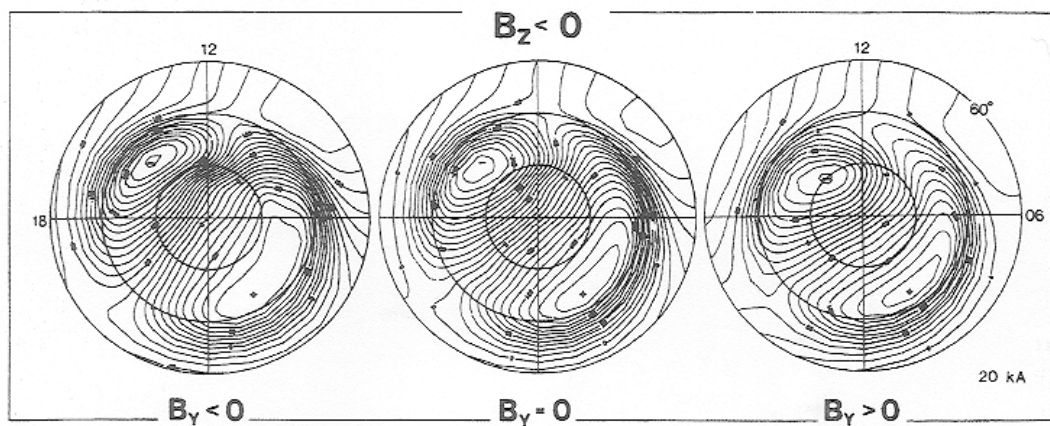
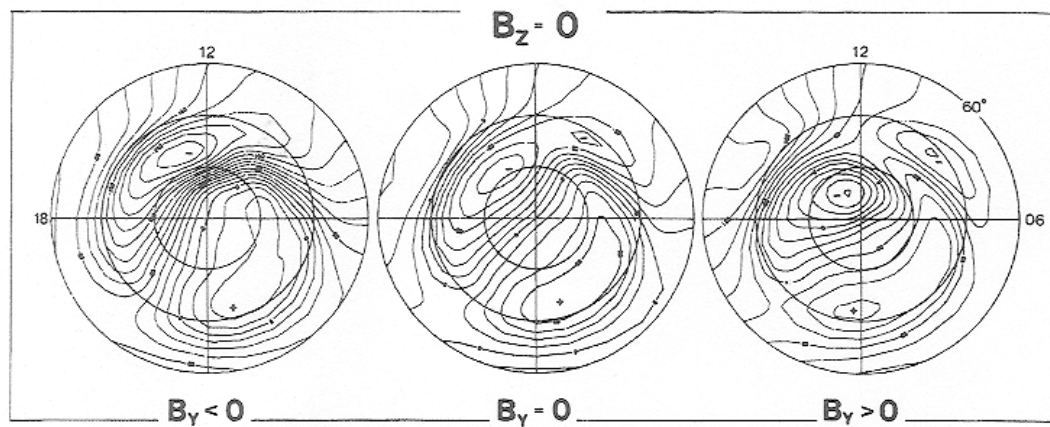
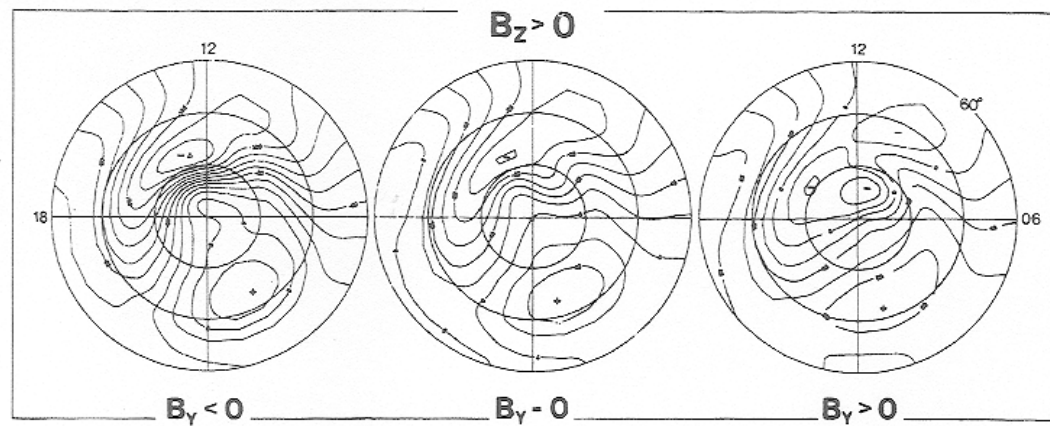


Figure from Menvielle and Marchaudon (2006)

"Geomagnetic Indices in Solar-Terrestrial Physics and Space Weather"

PC-North and PC-South
quantitative dimensionless indices



Equivalent ionospheric currents for various combinations of the interplanetary magnetic field components B_z and B_y

Figure from Friis-Christensen (1985)

Details of the PC indices

Objective:

monitor the cross polar cap ionospheric current

Contributors:

two observatories near the geomagnetic north and south poles

Physical quantity:

deviation of the horizontal magnetic field component from reference level

Reference level:

has been defined in various ways – unified definition under development

Method of derivation:

more complicated, several times revised – unified method in preparation

solar wind merging electric field

$$E_m = v_{sw} (B_y^2 + B_z^2)^{1/2} \sin^2(\theta/2)$$

transpolar current magnetic field

$$\Delta F_{PC} = \Delta H \sin \gamma \pm \Delta D \cos \gamma$$

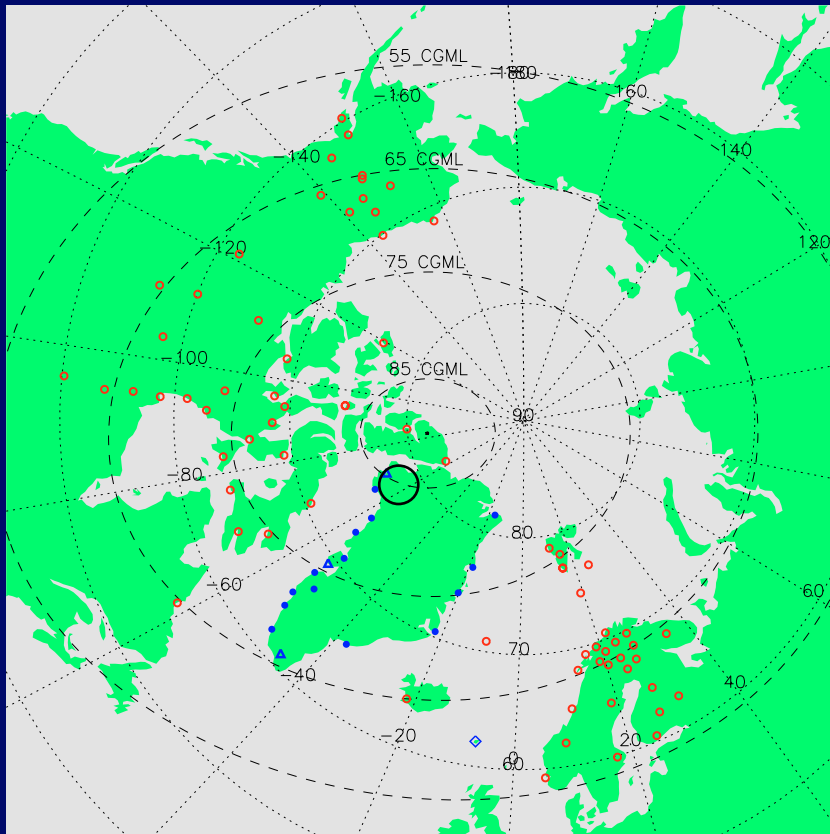
good correlation ($r > 0.8$)

$$\Delta F_{PC} = \alpha E_m + \beta$$

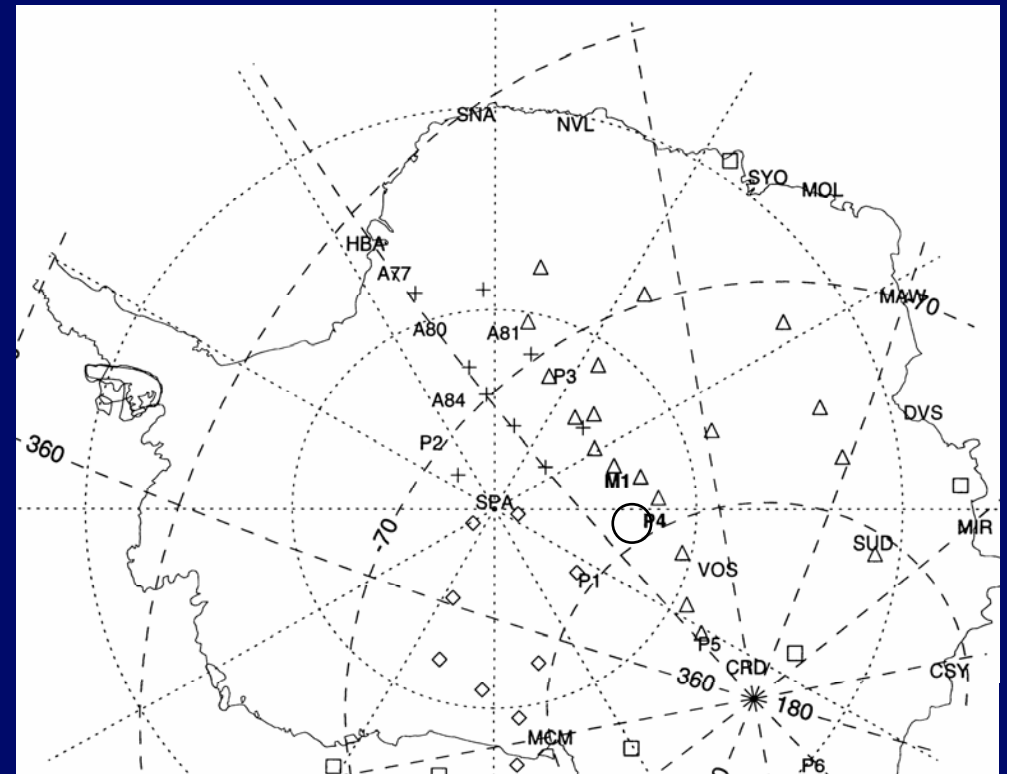
Characteristics:

represents in a statistically optimal way the solar wind merging electric field

Qaanaaq observatory (PC North)



Vostok observatory (PC South)



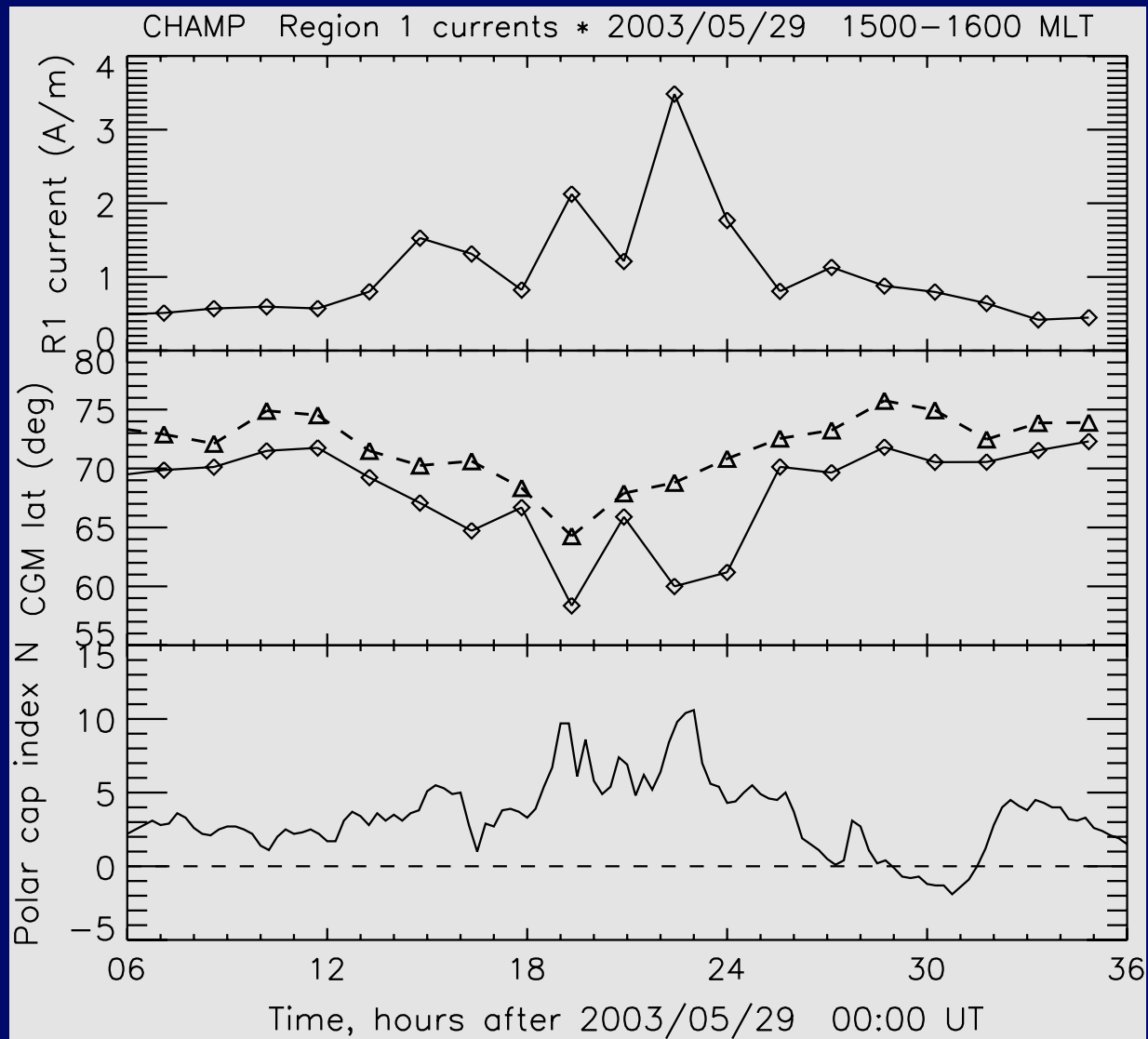


Figure from Hanuise et al. (2006)

Indices: Summary

- (1) SYM-H and ASY-H are possibly the most appropriate storm indices for space weather applications. Being relatively recent they are less suitable for long-term studies.
- (2) The Kp group of indices is generally not very suitable for space weather purposes although Kp is used a lot – unfortunately.
- (3) The high-latitude indices (AE group and PC group) are useful for both space weather research and operational applications