



**The Abdus Salam
International Centre for Theoretical Physics**



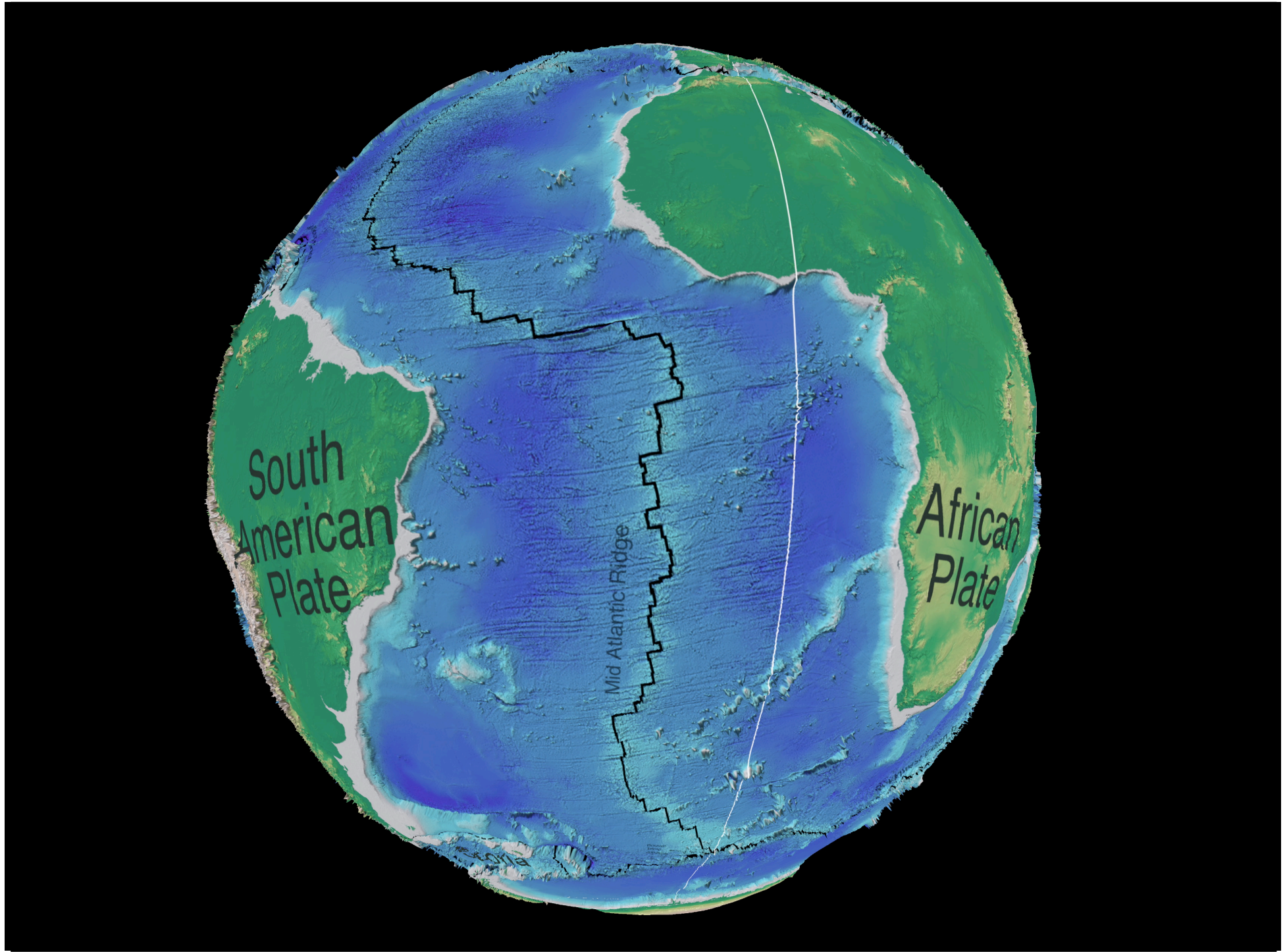
2053-10

**Advanced Workshop on Evaluating, Monitoring and Communicating
Volcanic and Seismic Hazards in East Africa**

17 - 28 August 2009

**Time and length scales of deformation in the East African Rift: Implications for EQ
and Volcanic hazards**

Cindy Ebinger
*University of Rochester
USA*



South
American
Plate

Mid Atlantic Ridge

African
Plate

Time and length scales of faulting and magmatism within the East African rift system

Cindy Ebinger, University of Rochester
Work in progress from Manahloh
Belachew and Dustin Coté



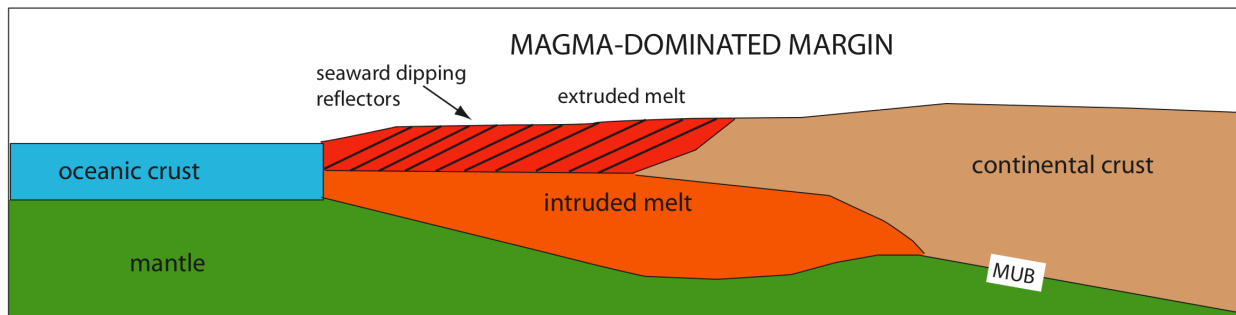
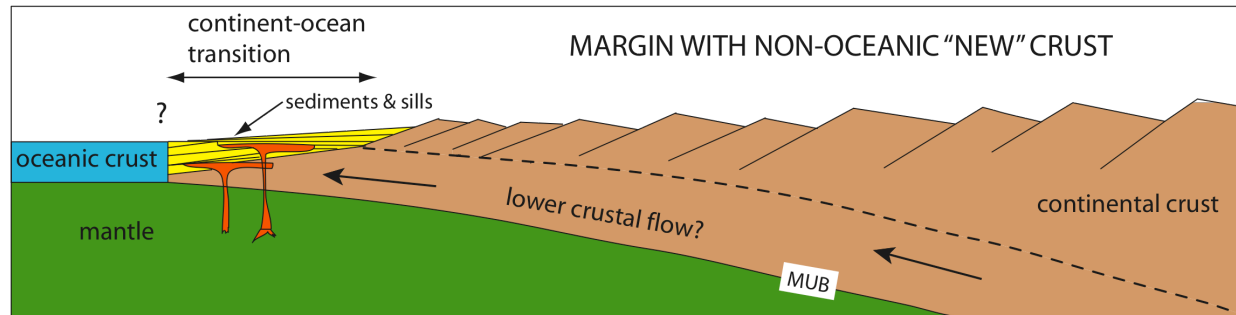
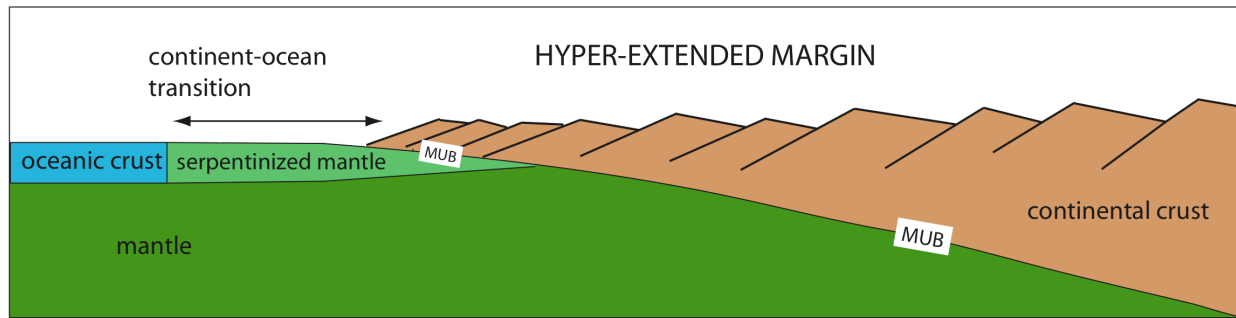
UNIVERSITY of
ROCHESTER

Addis Ababa
University
(Since 1950)



NATURAL
ENVIRONMENT
RESEARCH COUNCIL

Current conceptual models for continental rapture



MUB = mafic-ultramafic boundary

TEAMWORK – MULTIDISCIPLINARY STUDIES



Geodesicists, structural geologists, petrologists, seismologists

Technical, software, hardware support from PASSCAL/Seis-UK, UNAVCO, AAU, UDSM, UCLAS, etc



HARD WORK



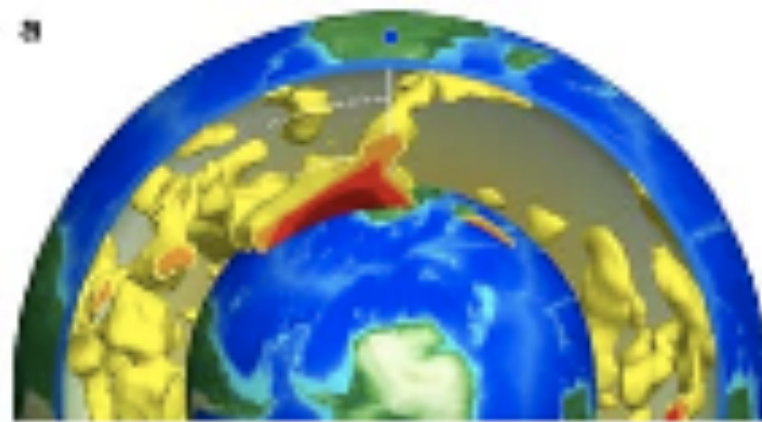
**Science
Communication**



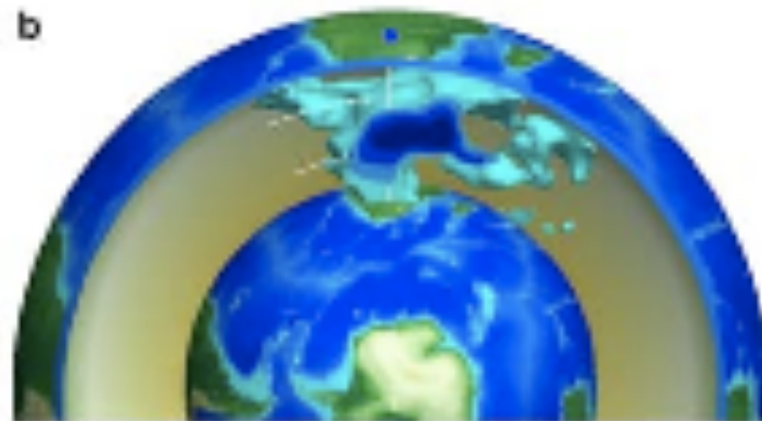
Implications for Rates

- Slow spreading - 3 -6 mm/yr in EAR south of triple junction; 20 mm/yr north
- During rifting crises, rates of ~m/day possible

- Implications - High instantaneous rates, plus growing population centers in rift increases seismic and volcanic risk to that of some subduction settings



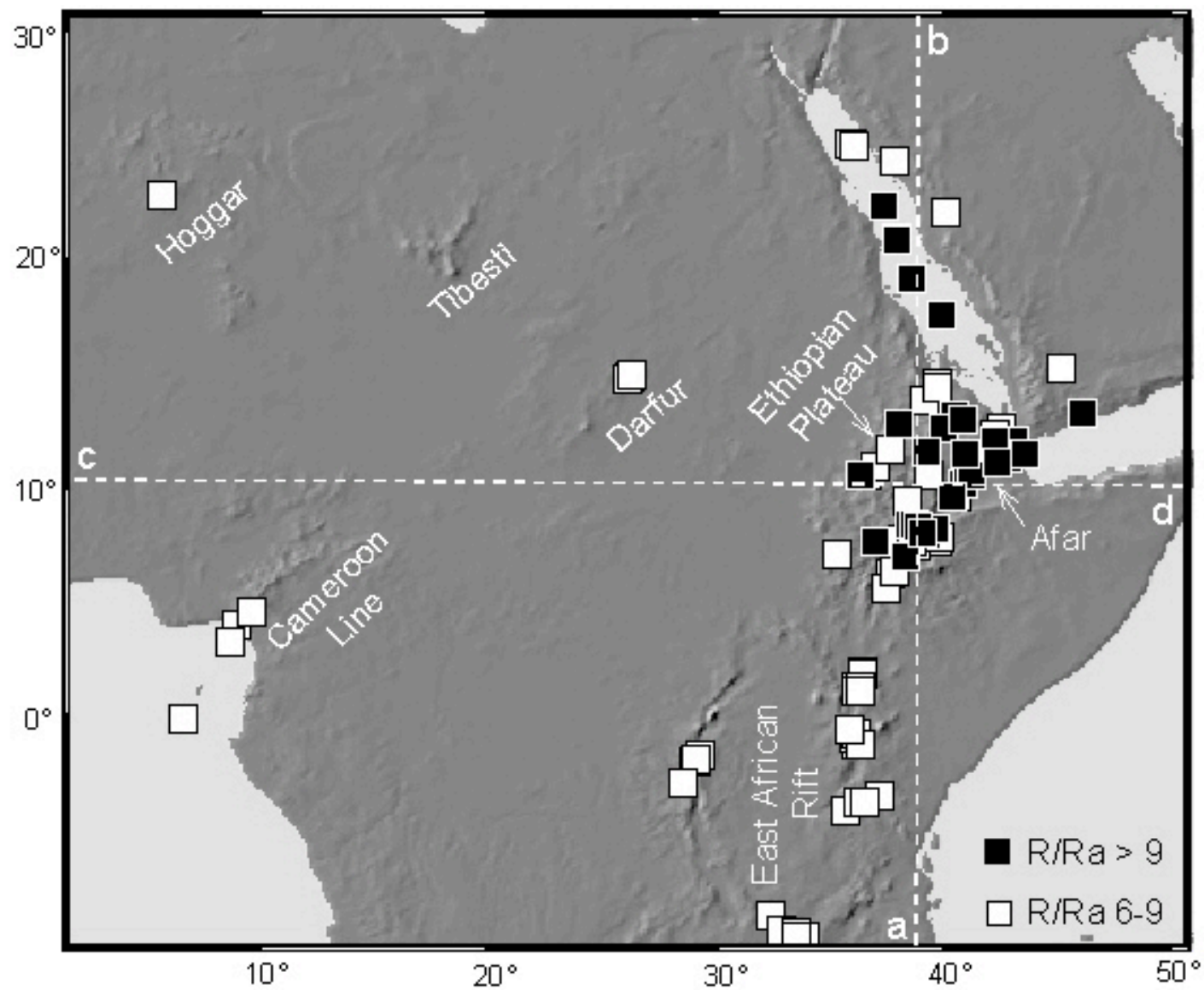
-0.13 -0.15 -0.20 -0.30
 $\delta\rho - \text{thermal} + \text{chemical} (\%)$



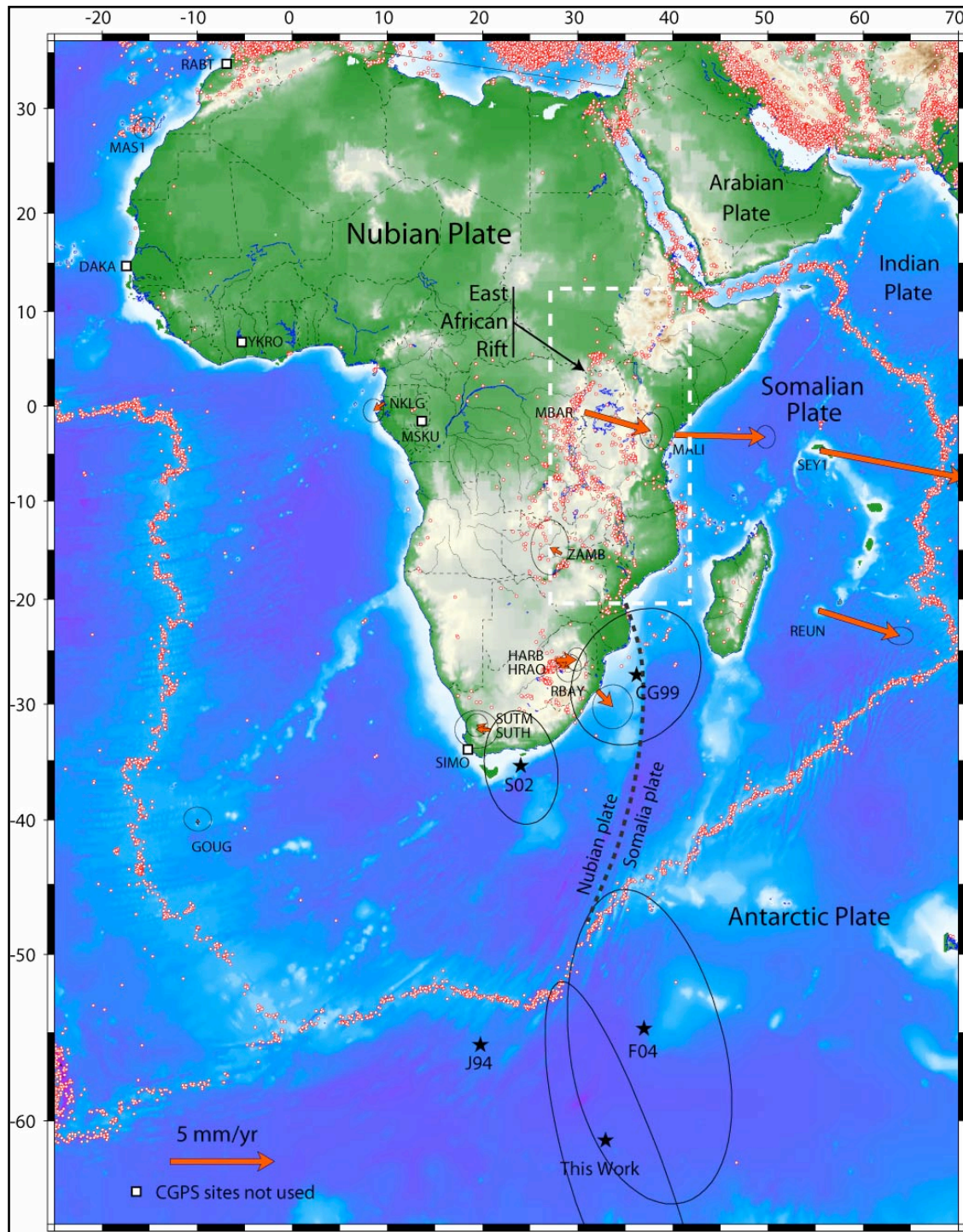
0.30 0.35 0.45 0.55
 B - buoyancy ratio

Figure 4. Total density perturbation field and buoyancy ratio $B = |\delta\rho^{\text{chemical}}/\delta\rho^{\text{thermal}}|$ within the African superplume region. (a) Summation of the thermally- and chemically-induced density fields yields negative density (positive buoyancy) of the superplume structure. (b) The buoyancy ratio field reveals that the superplume buoyancy is strongly reduced by the positive-density chemical component near 1800 km depth. The superplume structure appears to bend eastward within the same region as a response to buoyancy deficiency.

Simmons et al., 2007

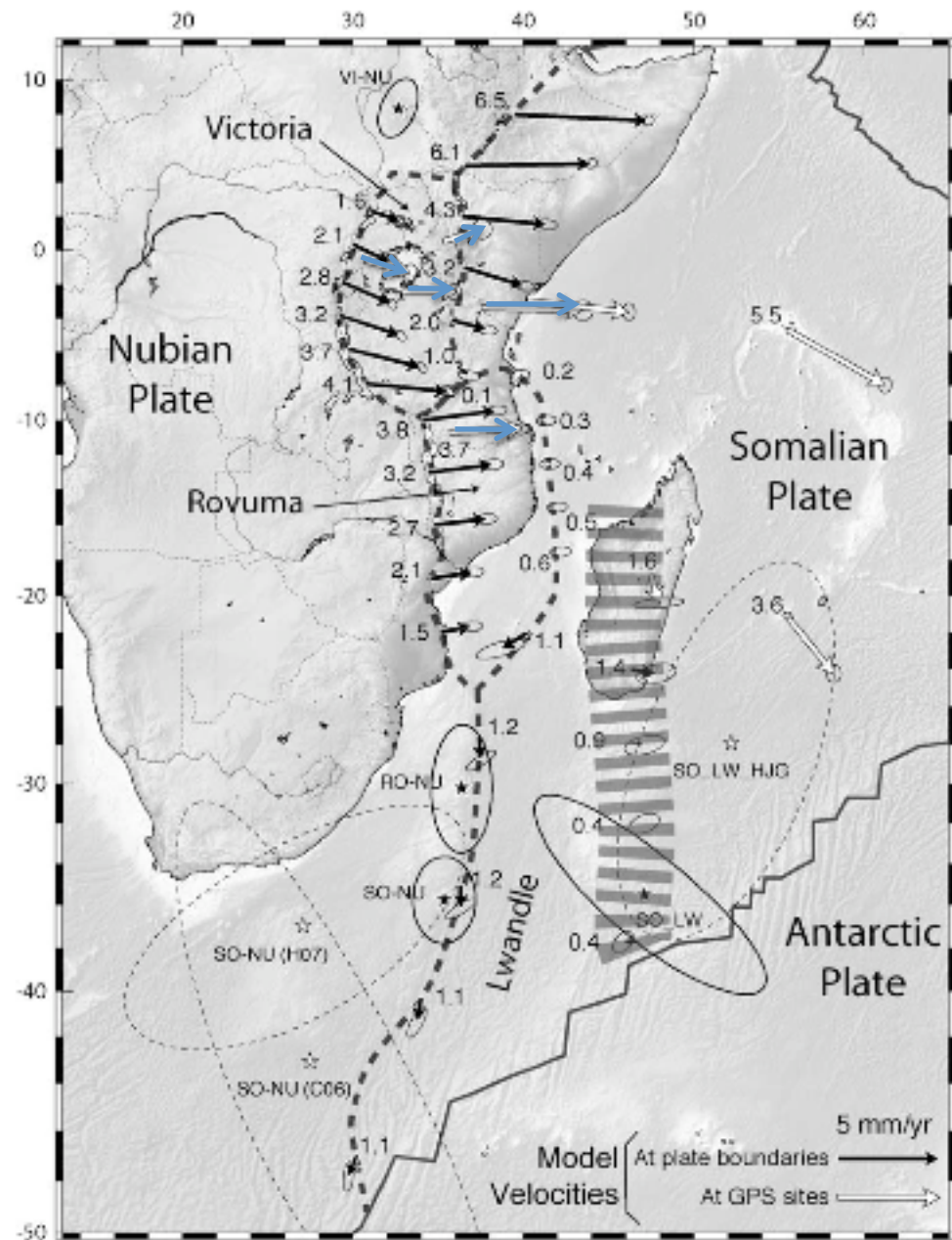
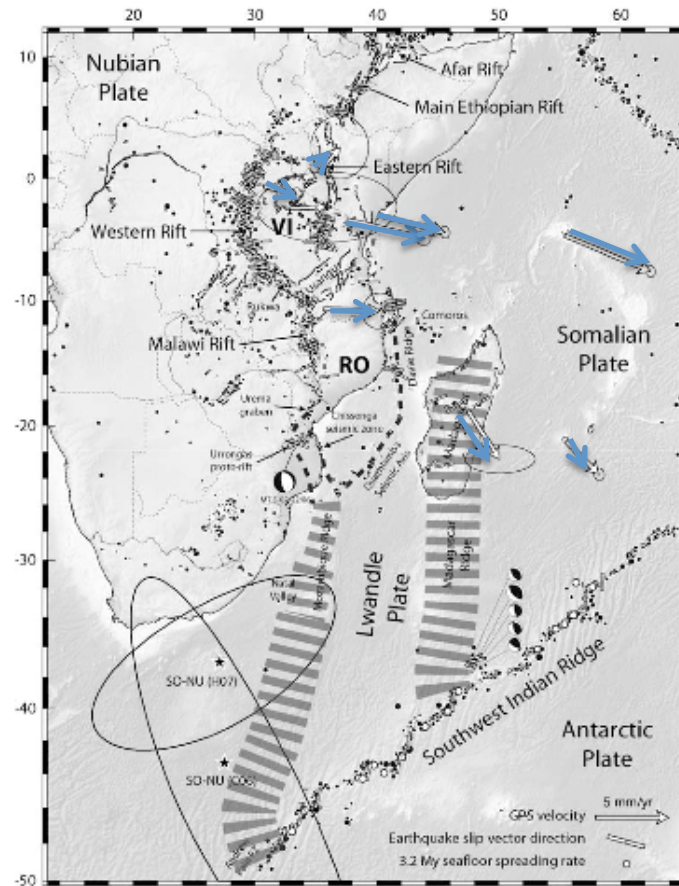


He anomaly patterns - Pik et al., 2006



Chris Hartnady noted a differential opening direction between WR and ER, and suggested that central Tanzania Archaean craton was a discrete microplate.

Fig. from Calais et al., 2006



Stamps, Calais et al. GRL 08

Time Scales of Rifting Cycle

- Far-field + body forces - decades to centuries - depends on inherent strength, lateral density contrasts, plate pull
- Fault slip, magma intrusion - minutes to months
- Viscous relaxation - years
- Isostatic compensation for the deformation - 10^4 years
- Aseismic slip - ?????? need more collocated GPS + seismic

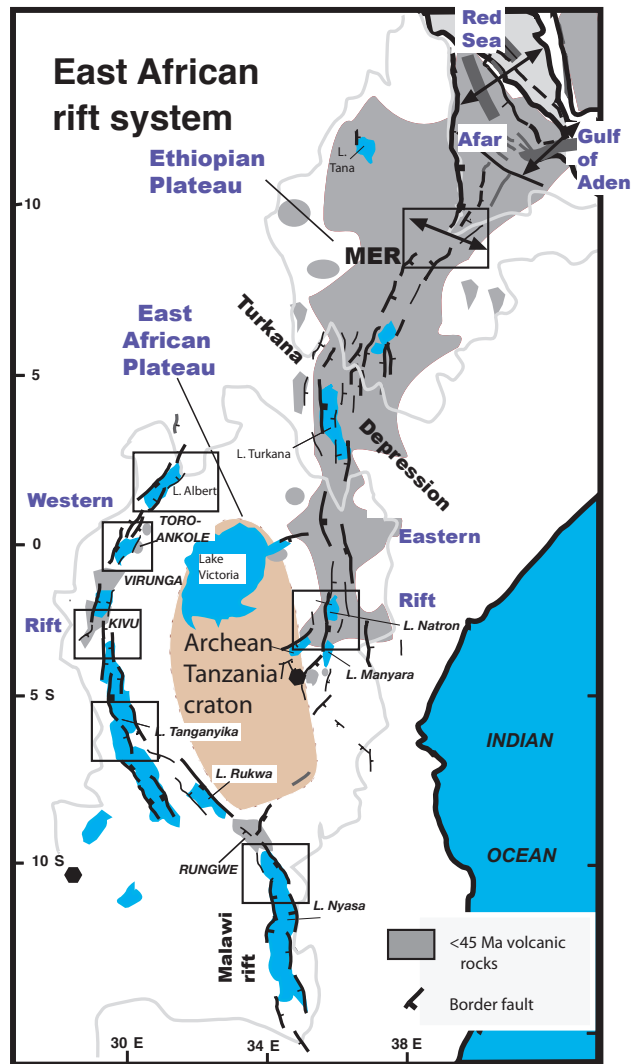
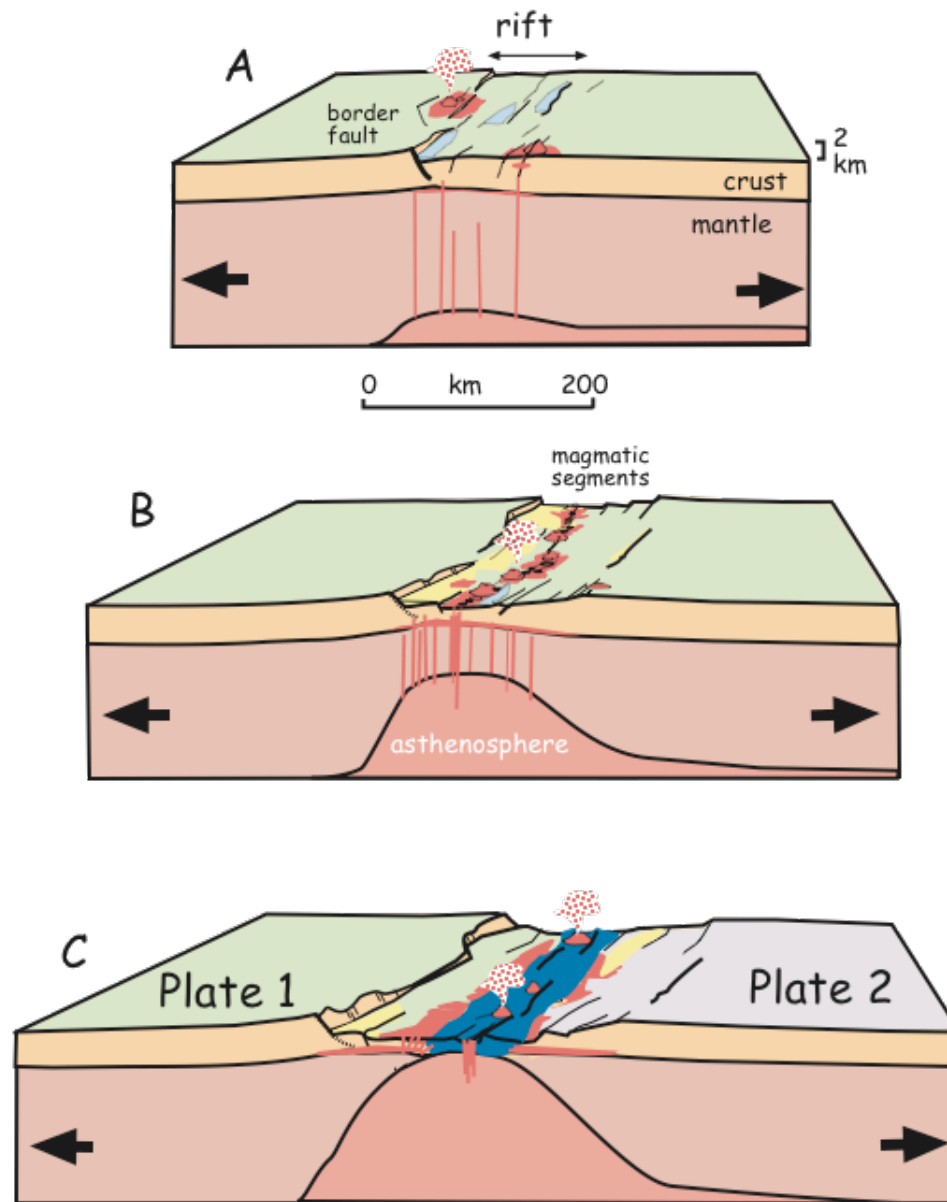


Fig. 5, Ebinger and Scholz



Rifting with Magma

Stage 1 - Thick lithosphere, long repeat times; large stored stresses.

Stage 2: Strain localization to zones of magma intrusion rather than older border faults

Stage 3: Onset of seafloor spreading similar to ultra-slow ridges - ridge jumps, 'off-axis' volcanism, and long repeat times.

Fig 3.

Consistent patterns - predictive model for rift basins worldwide

- *Interactions between magma and segmented large offset faults.*
- *How do they interact in space and time?*

Shillington talk later...

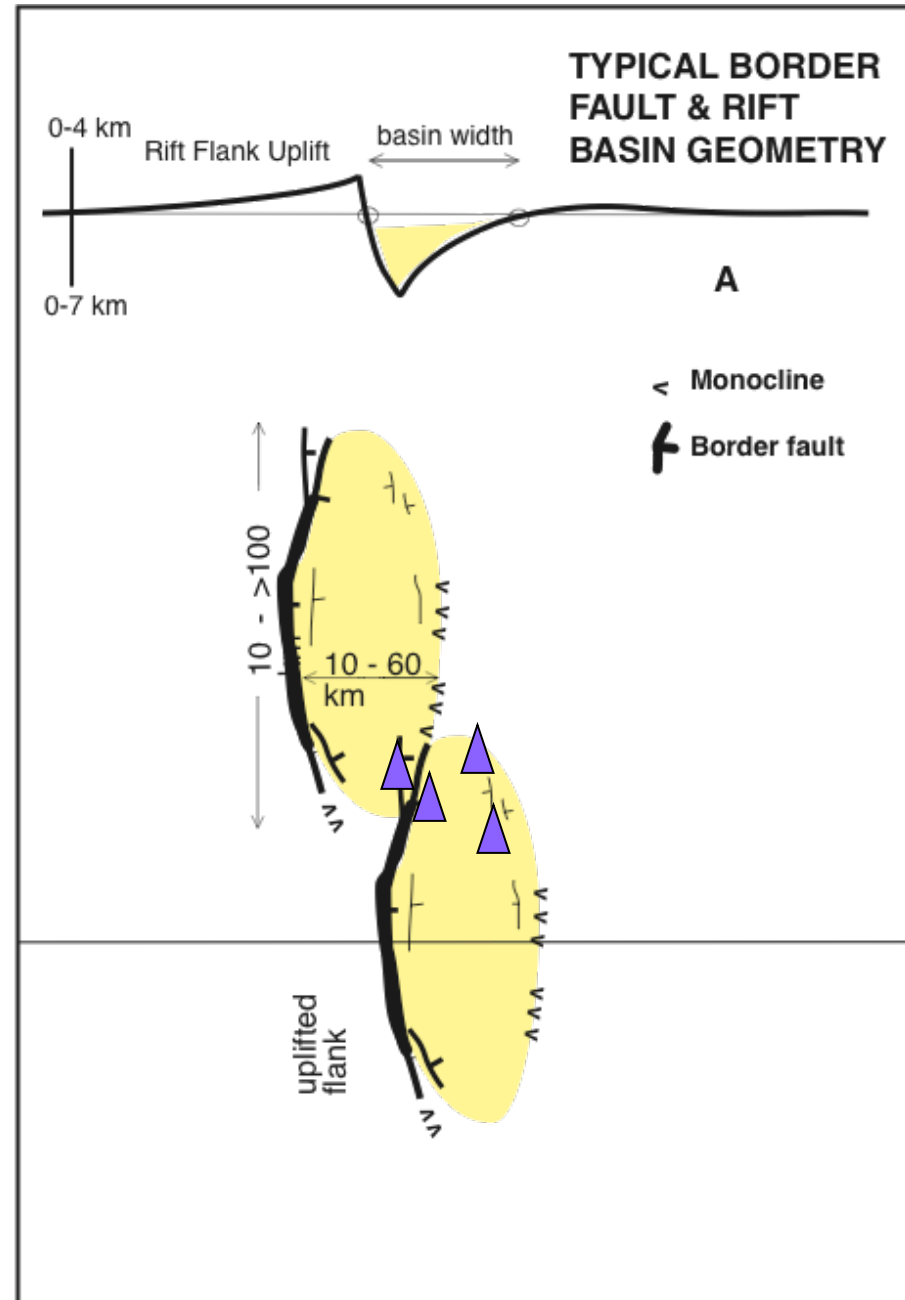
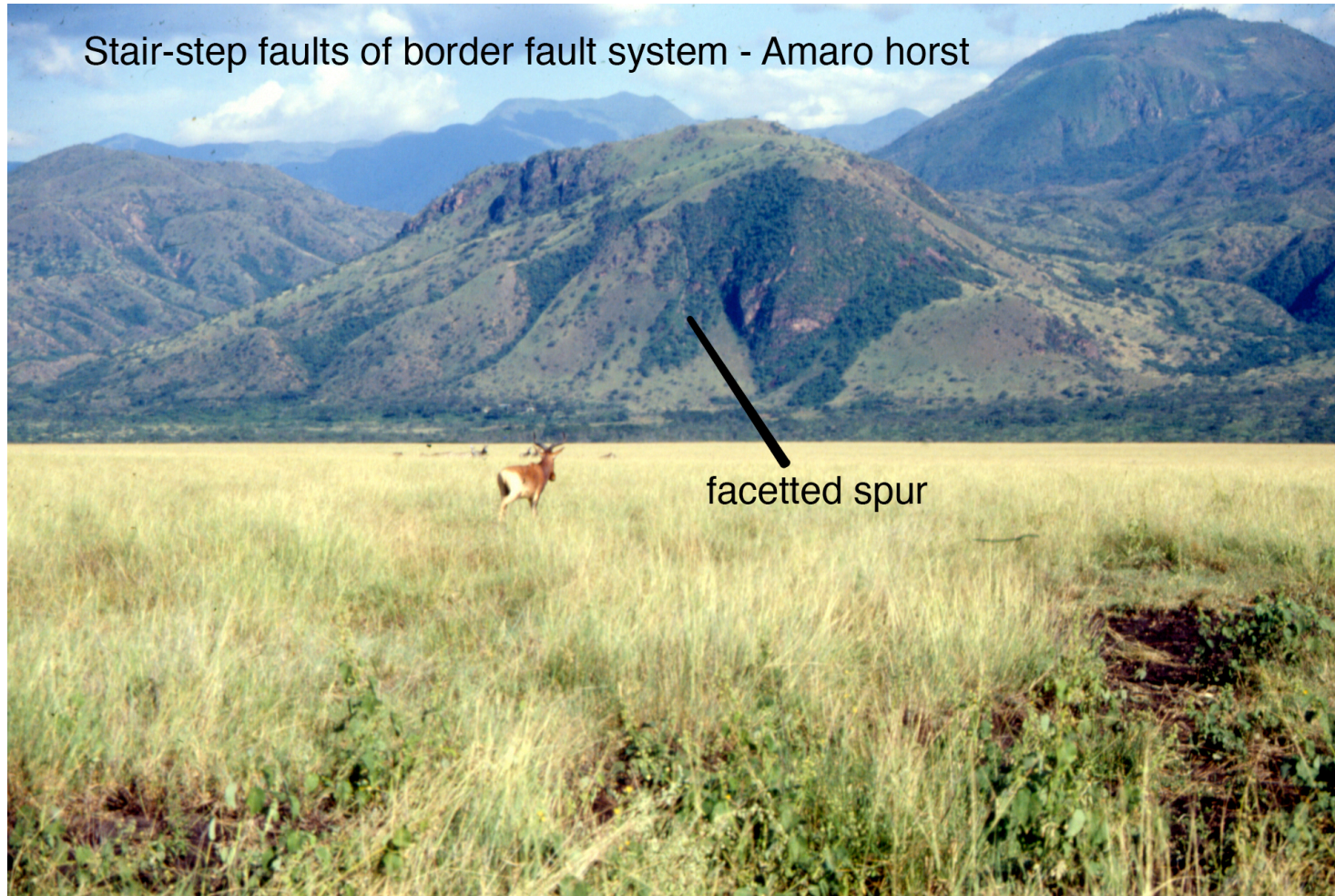


Figure 2. Ebinger et al.

Stair-step faults of border fault system - Amaro horst

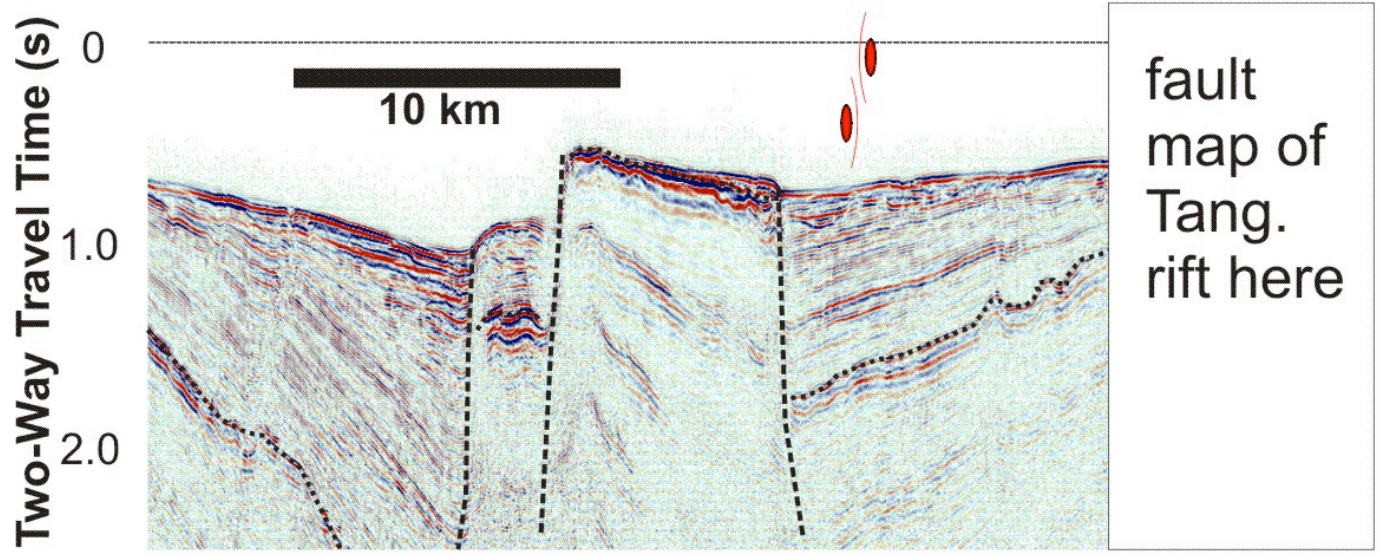
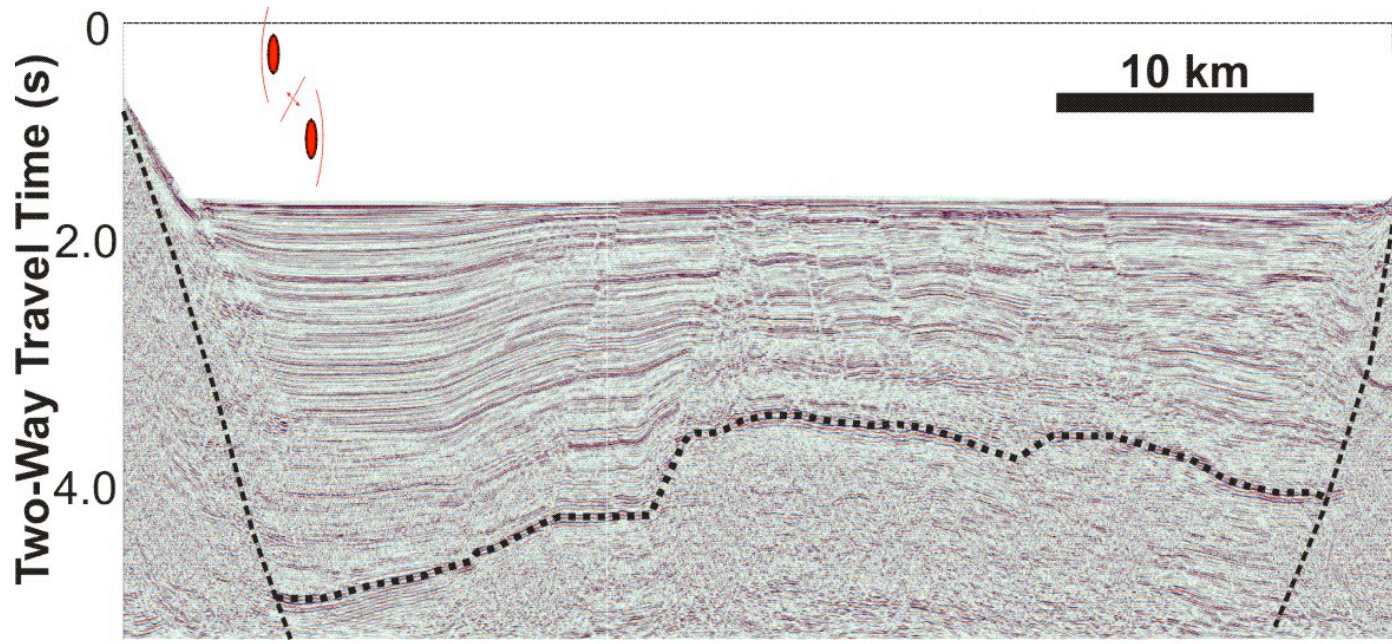


faceted spur

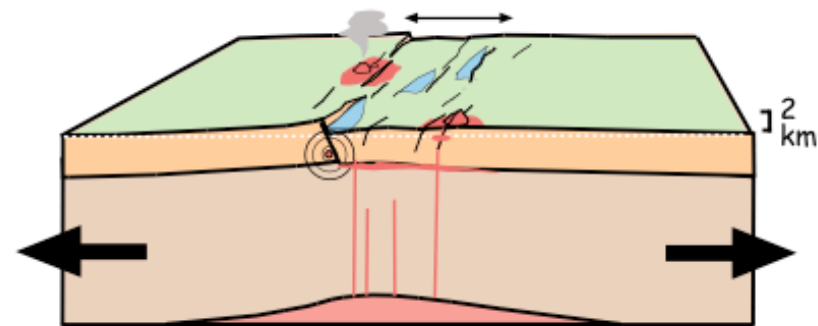
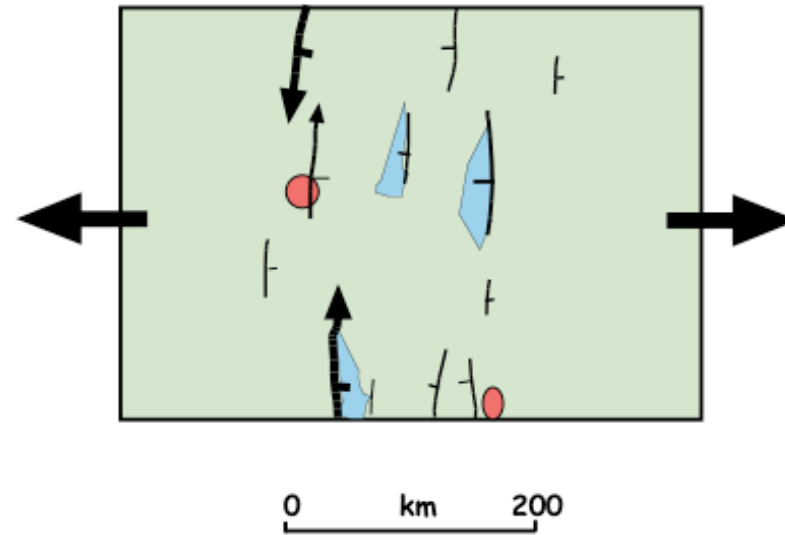
Kivu basin

Accommodation zone





Ebinger and Scholz, in Busby and Azor, Sedimentary Basins



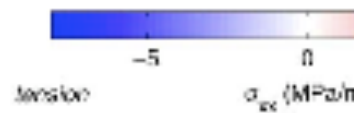
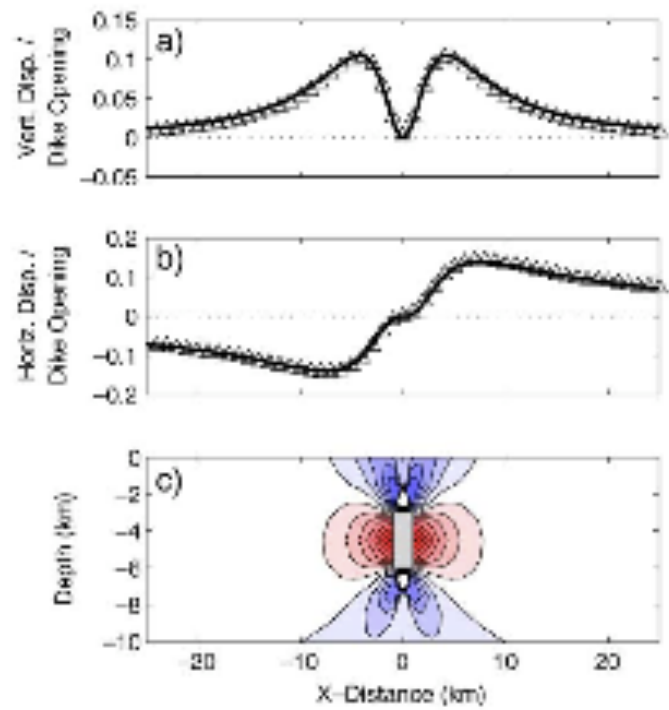
Front face ~ KRISP 94

Natron basin event, 2007 d'Oreye et al., Albaric et al.

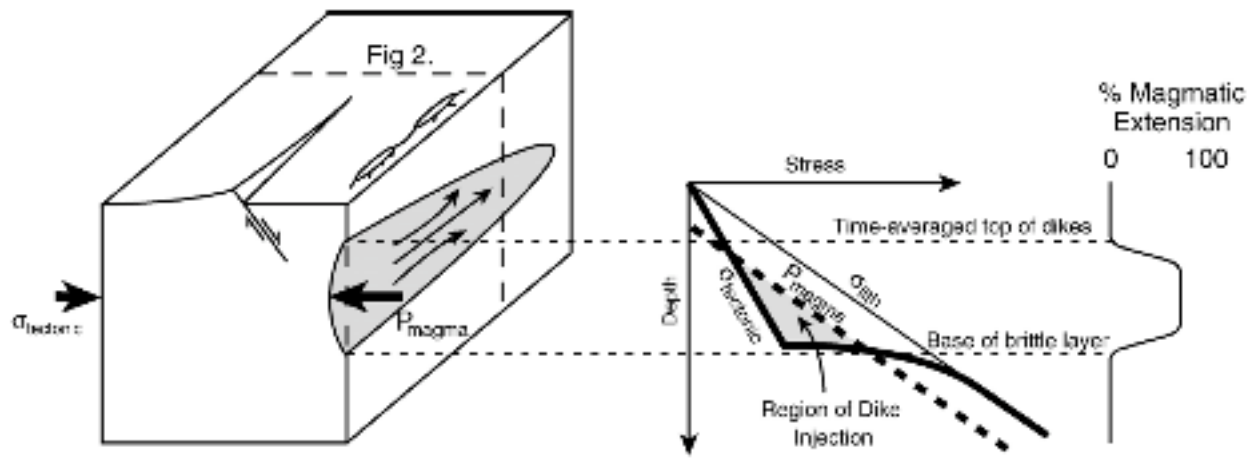


October 19 eruption of Ol
Doinyo Lengai

Photo by Majura Songo,
UDSM



Behn, Buck, Sacks,
EPSL, 2006

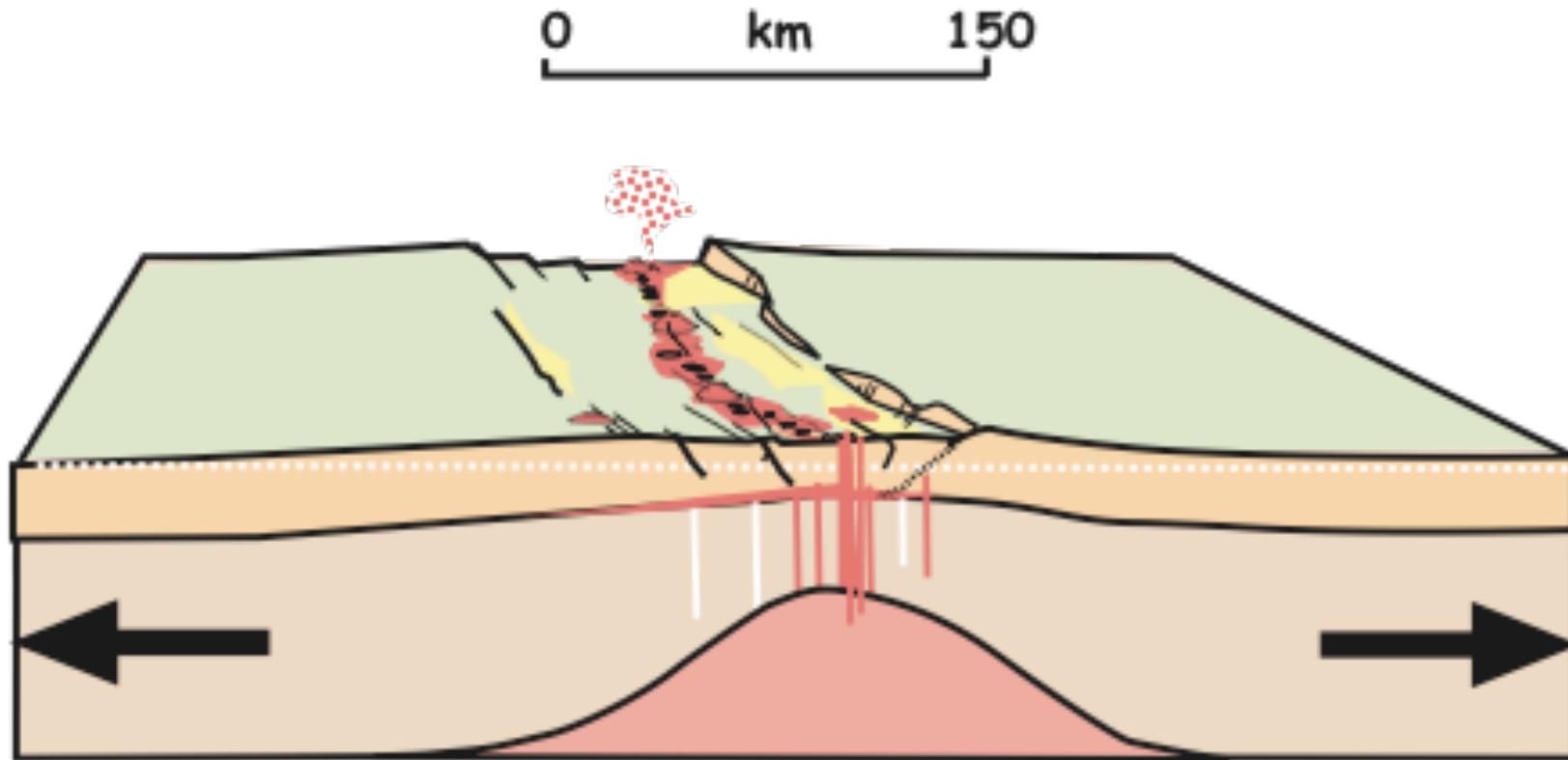


Implications for Hazards

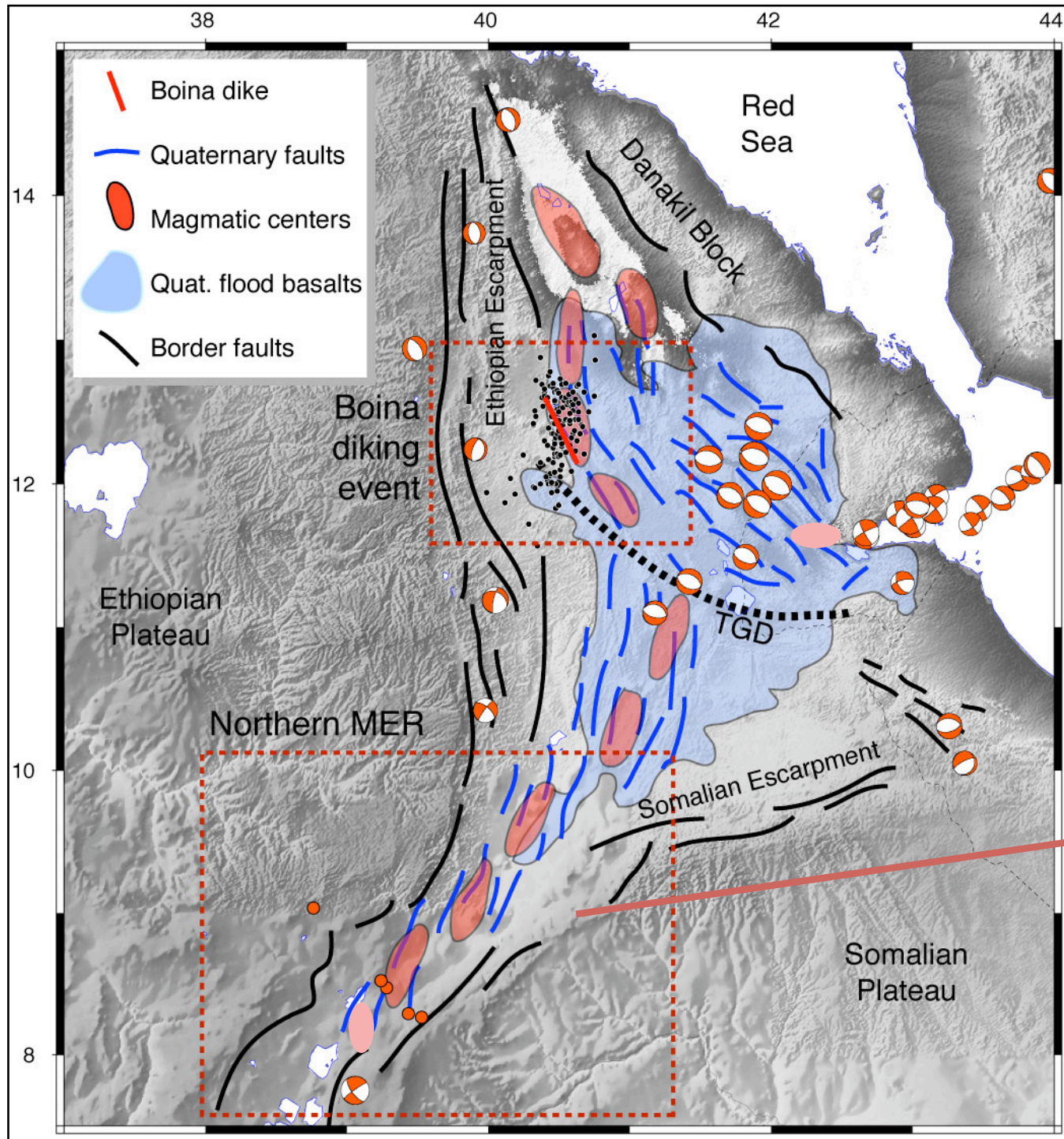
- Magma intrusion via 'hidden' subsurface dikes - hidden process
- Rifting crises, rates of \sim m/day possible
- Implications - High instantaneous rates, plus growing population centers in rift increases seismic and volcanic risk to that of some subduction settings

Length Scales of Rifting Cycle

- Regional magma-plate boundary situation
1000 km + scale
- Along-axis segmentation in both faulting
and magmatism -50 to 100 km in EAR (may
decrease with time)
- Entire segment may open during single
episode - e.g., Dabbahu rifting episode;
Malawi rift - Jackson and Blenkinsop
example

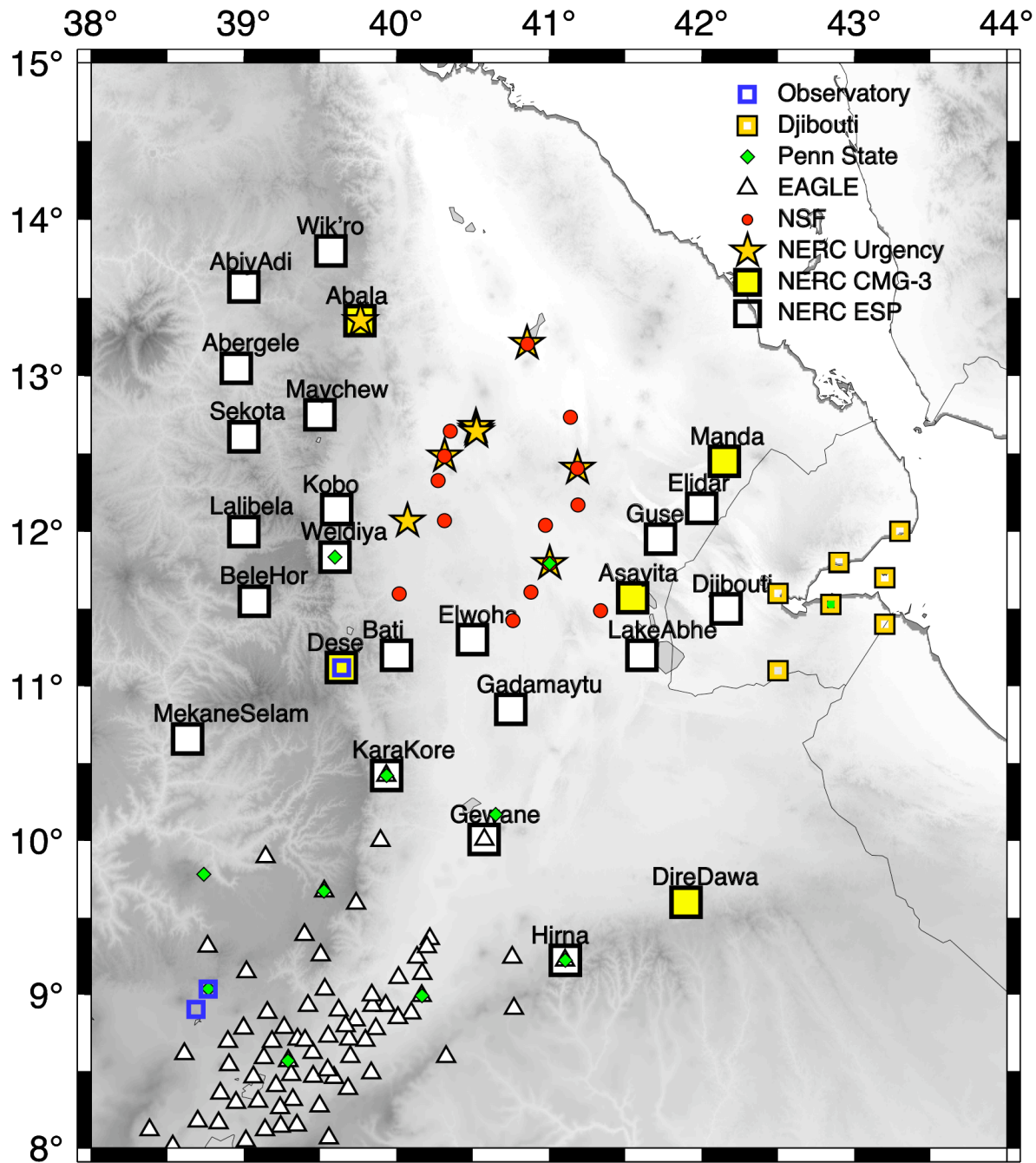


Minor crustal stretching, narrow zone of magma intrusion above mantle lithosphere thinned to <70 km = READY SUPPLY OF MELT

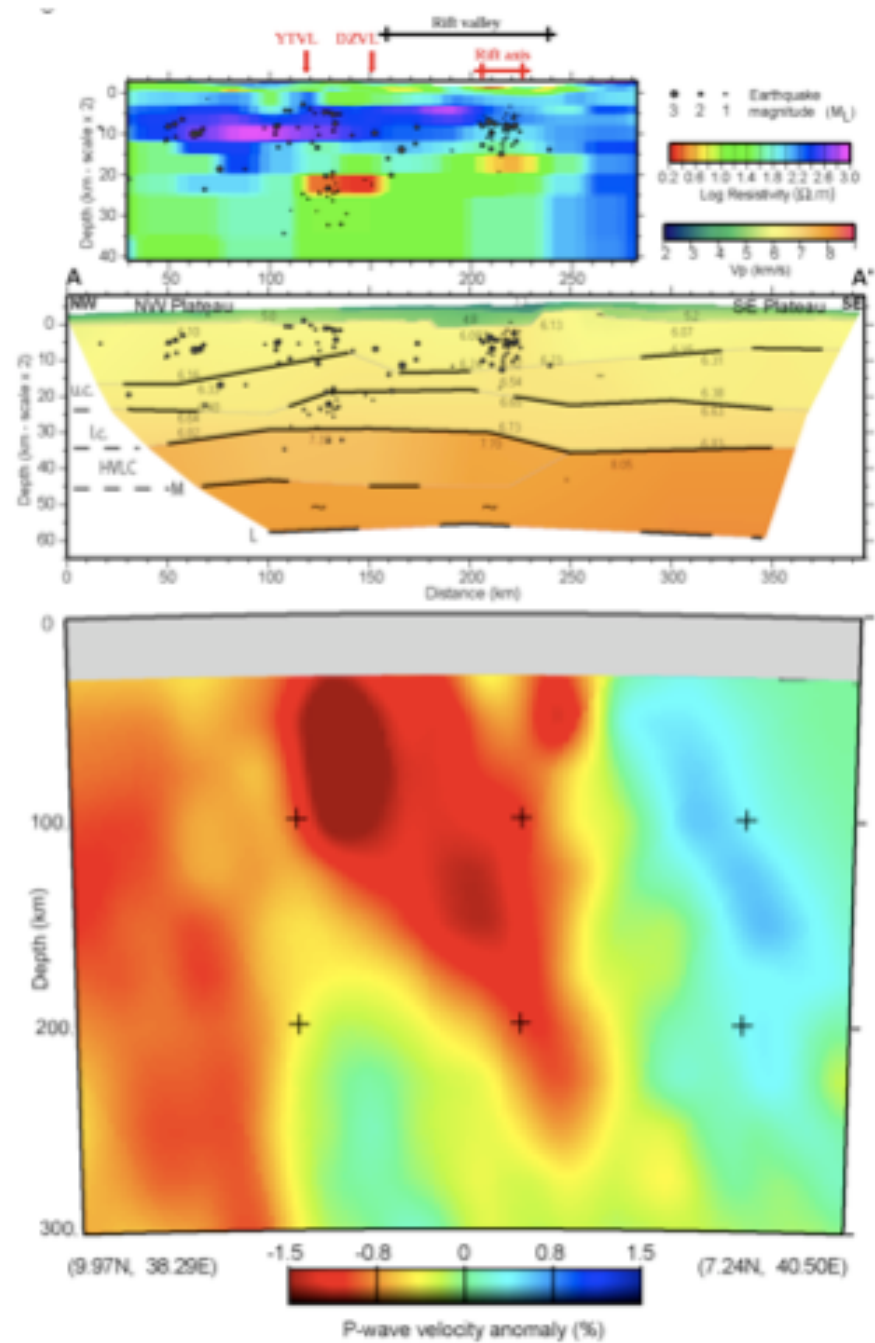


Quaternary strain localised to ~60 km long zones of fissures, aligned eruptive centres and faults - "magmatic segments"

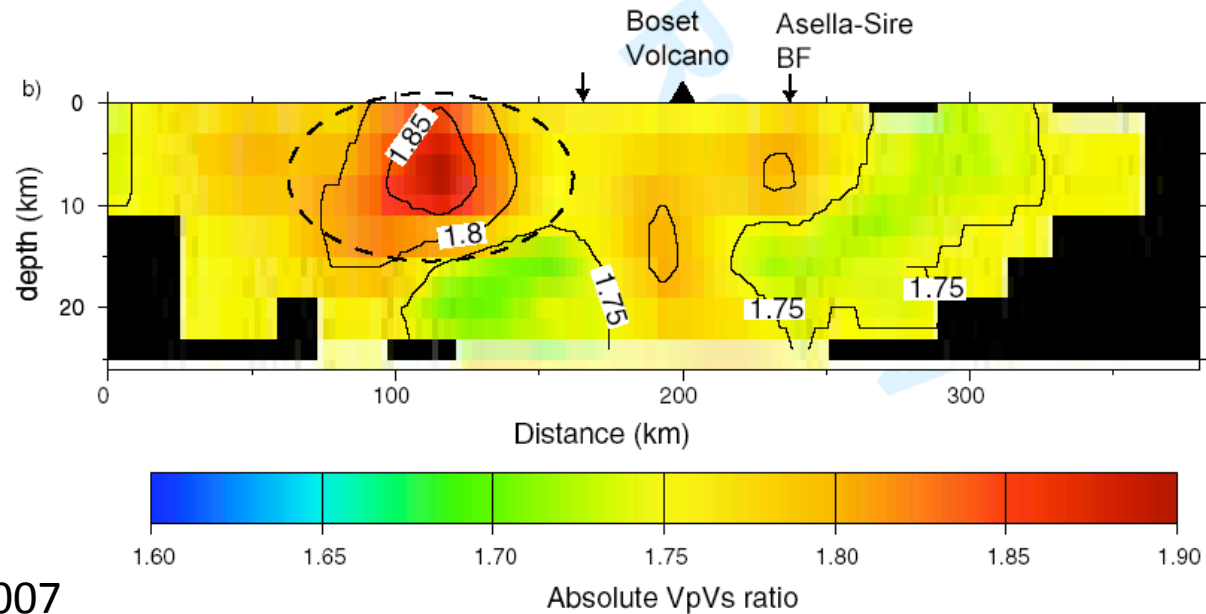
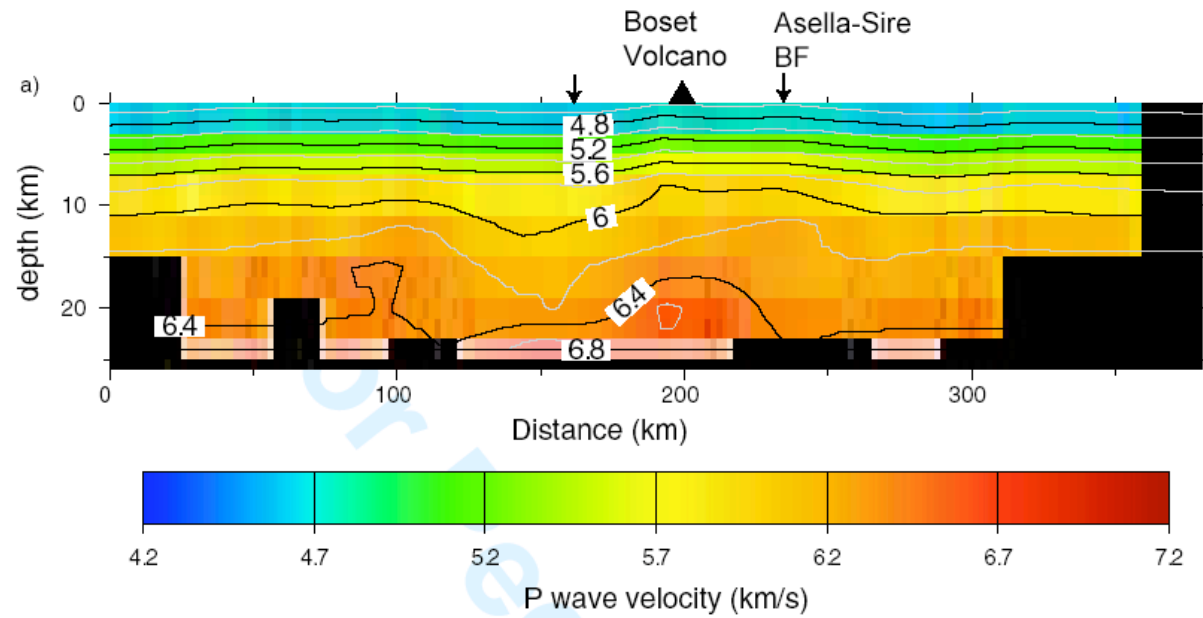
EAGLE probed structure beneath magmatic segments - underlain by mafic intrusions and LVZ in thin mantle

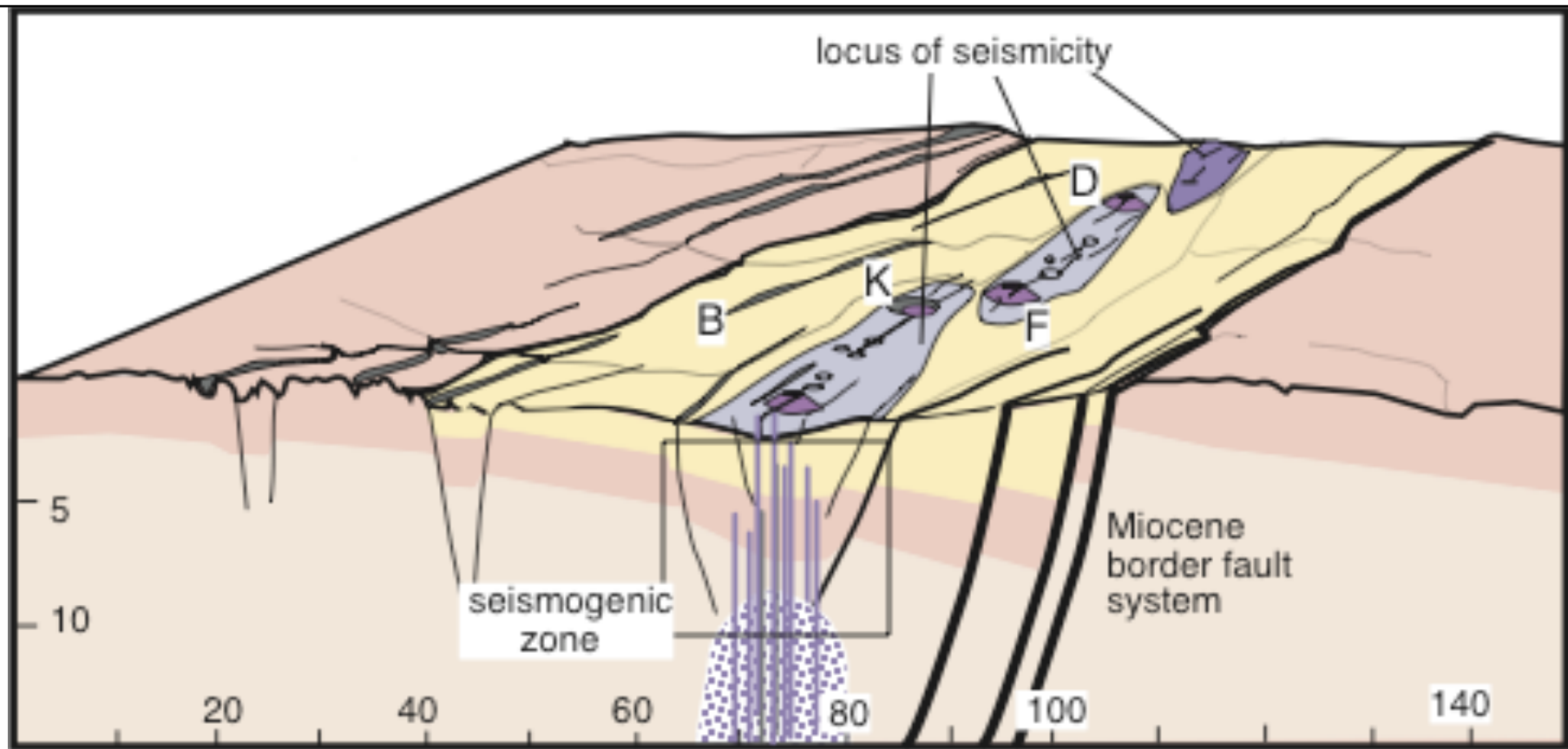


US-UK-Ethiopia
partnership;
Triangles –
Data available for
teaching, research;
Rest – work in
progress – France,
UK, US, Ethiopia,
Djibouti



Keir, Bastow,
Cornwell, Whaler
Gcubed, 2009





Crustal segmentation defined by narrow zones of mafic intrusions that accommodate strain below 10 km; seismicity triggered by dike intrusions and dike-induced faulting. Steep faults; >50% have eruptive products along their length (<10 km- Casey et al., 2006).

Keir et al., JGR, 2006

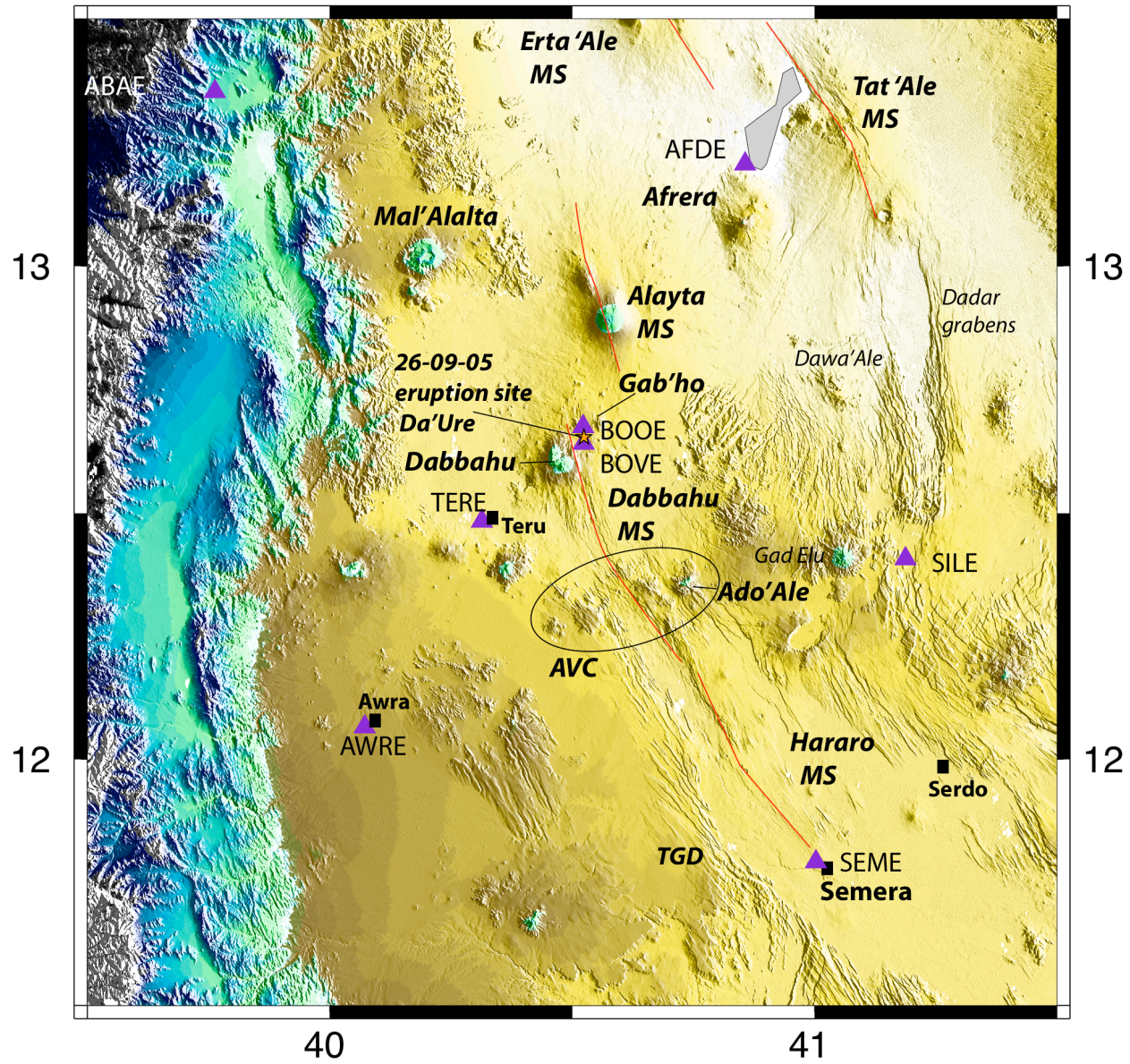


Figure 2. Ebinger et al. (revised)

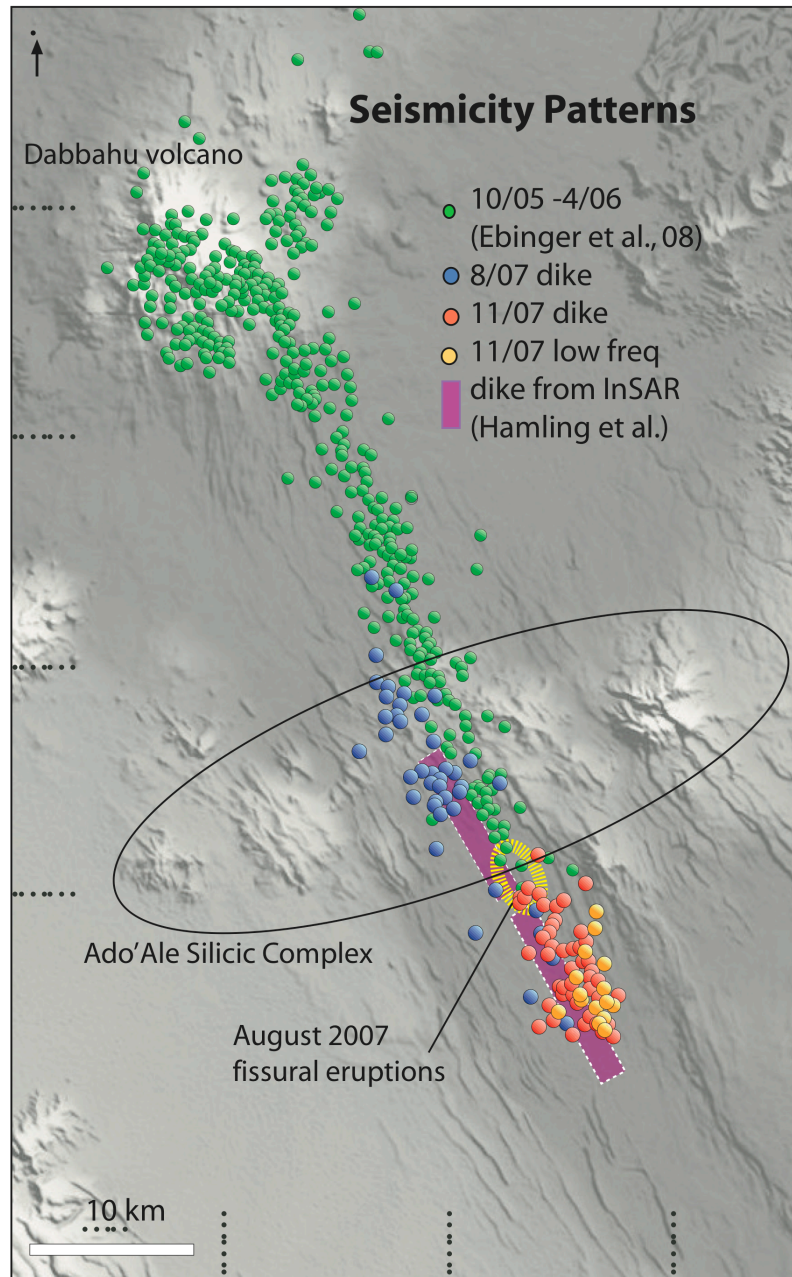


March expedition to mid-segment. JRowland

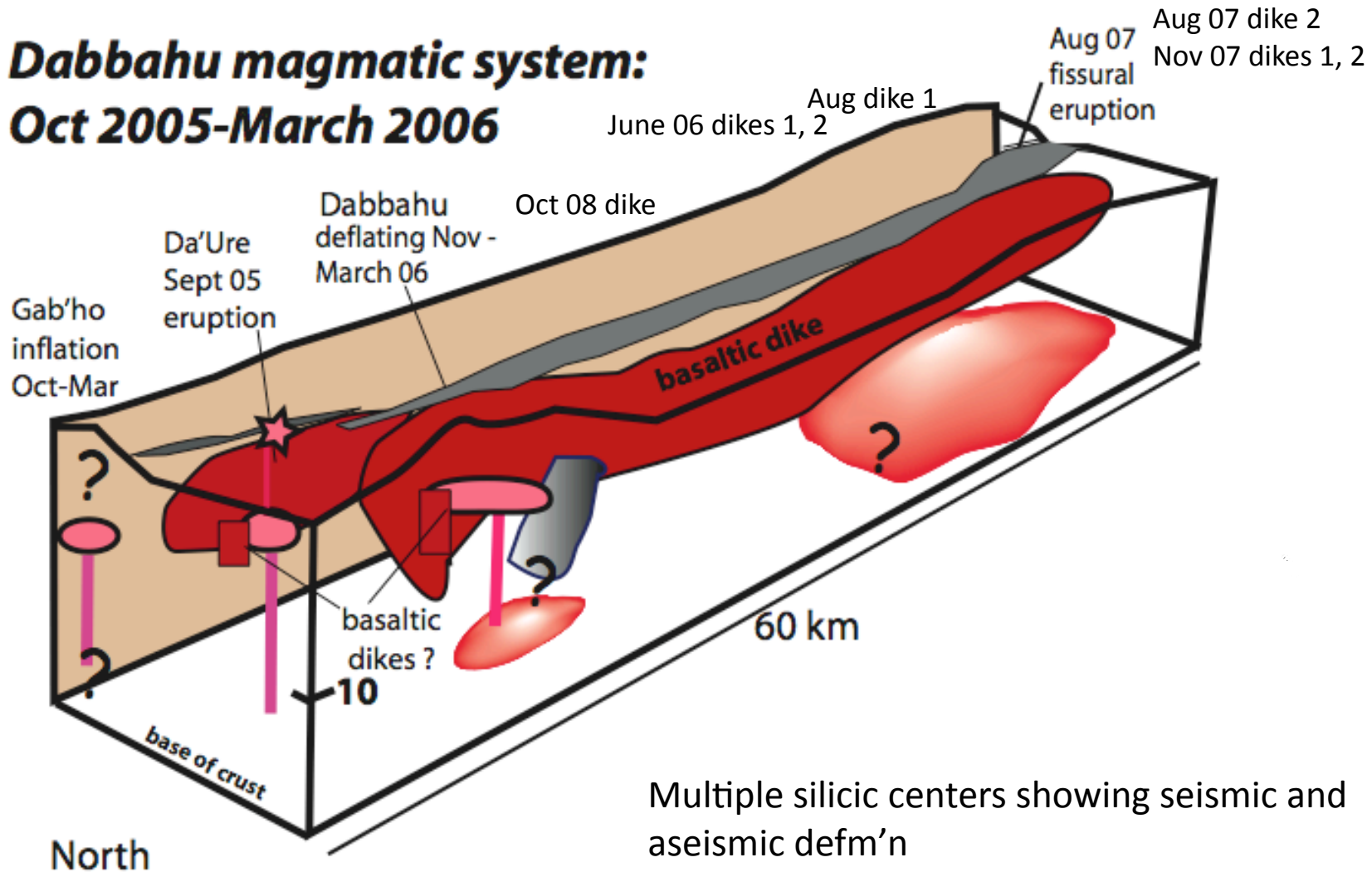


Liz Baker

3m-high white scarps - slip in July-Sept '06 - E Baker



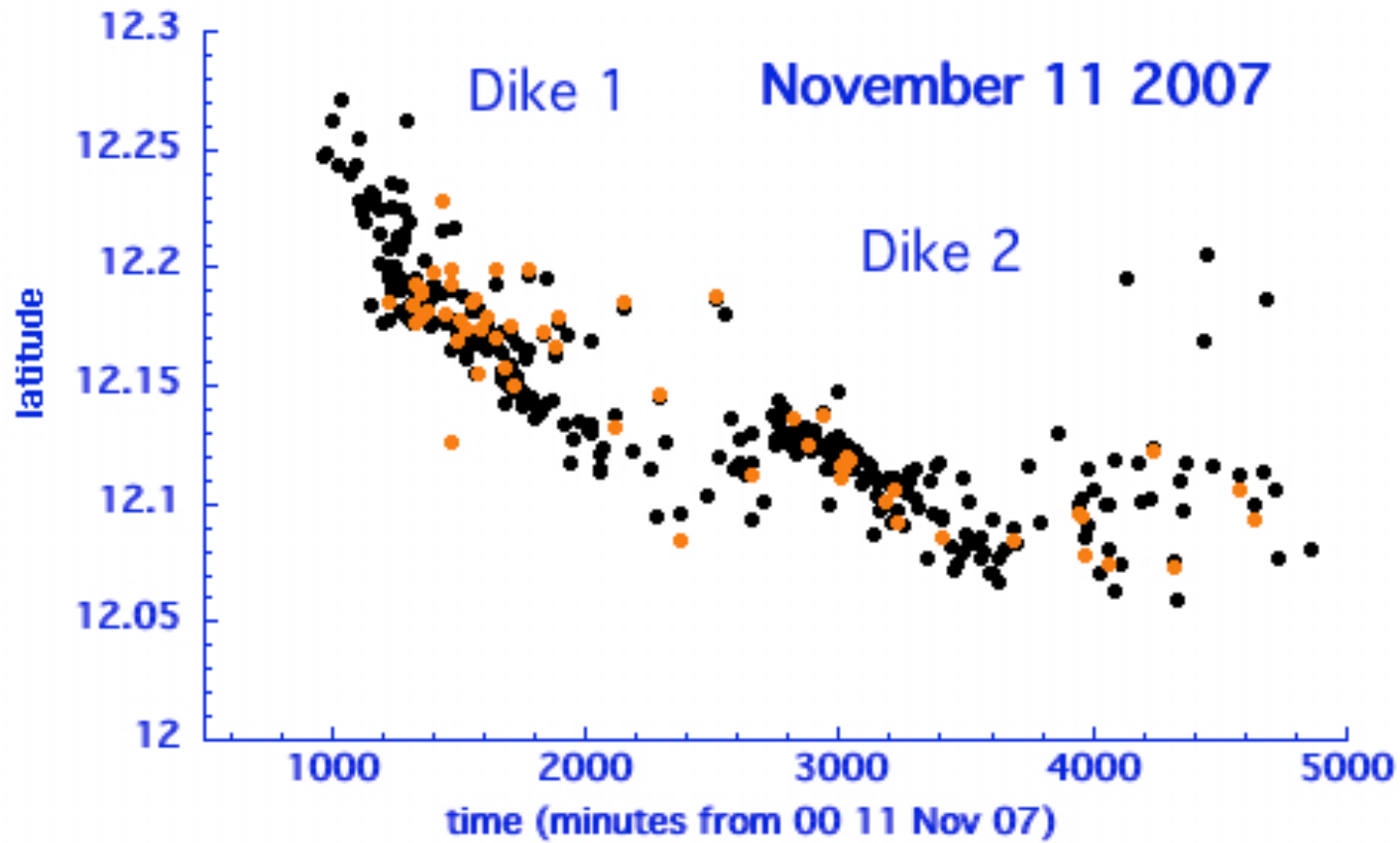
Dabbahu magmatic system: Oct 2005-March 2006

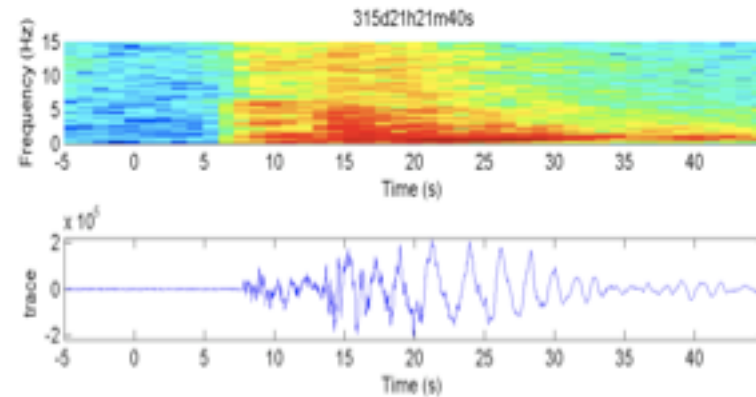
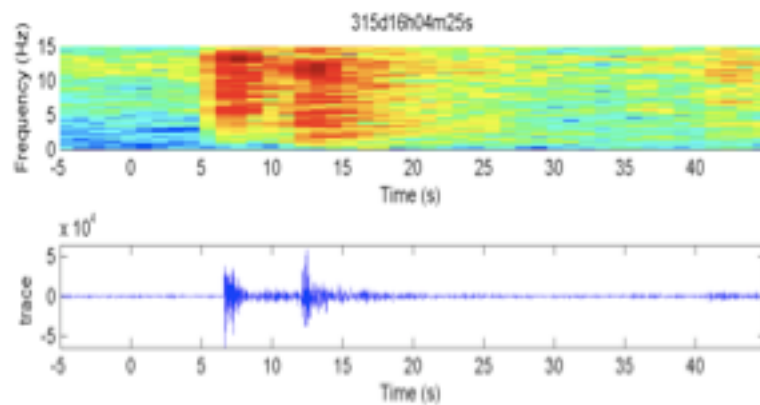


Multiple silicic centers showing seismic and aseismic defm'n
 No evidence for shallow chamber feeding dikes

- Centrally fed, ~60 km-long segment
- Dikes propagate northward and southward from segment center
- Shallow magma chambers with fractionated lavas at northern tip of segment where abuts 'cold' lithosphere (may have triggered rifting event- Ayele et al., 2009)
- Base crust/asthenospheric source zone from preliminary crustal tomography (no chamber shallower than 7 km)
- Time varying seismicity patterns reveal magma injection process

June 06, August 07, November 07 dikes are each multiple dikes – each 10-20 km





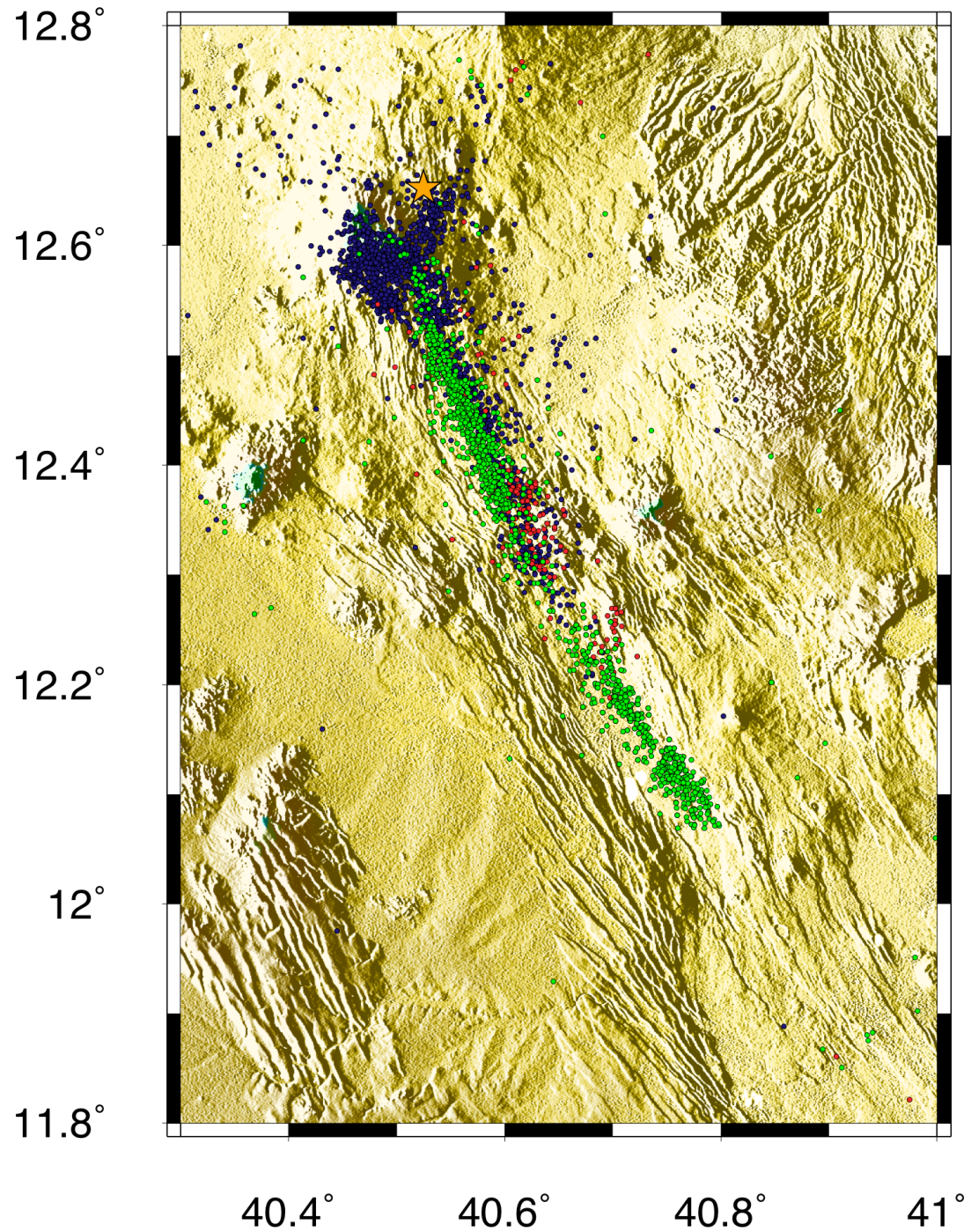
A comparison of a typical tectonic earthquake with a classic LF signal from the Nov 2007 swarm. The LF earthquake shows a clear 'P' onset that becomes superposed with a peak $< 0.5\text{ Hz}$ signal.

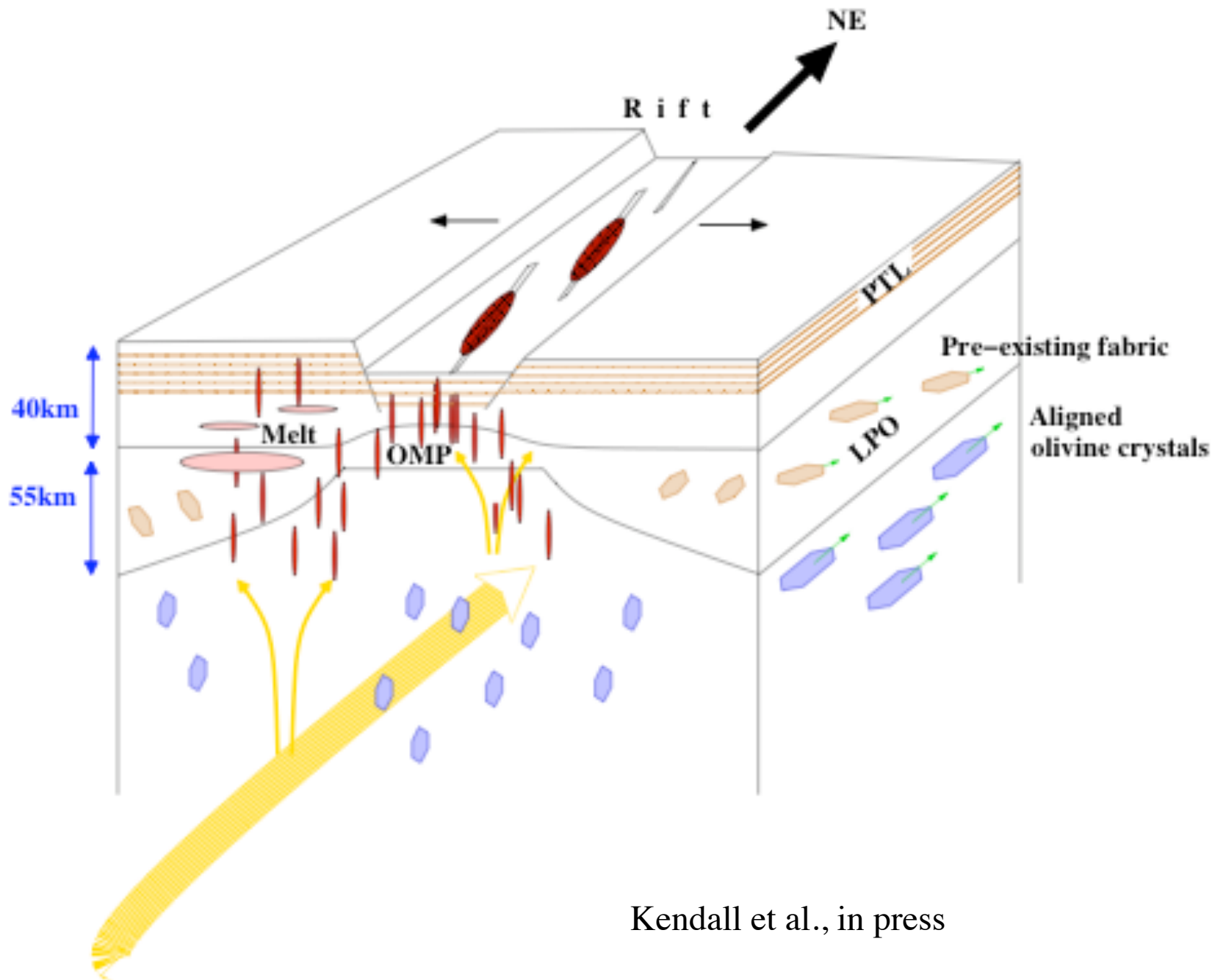
Dustin Cote, U Rochester

March 7 to
August '07

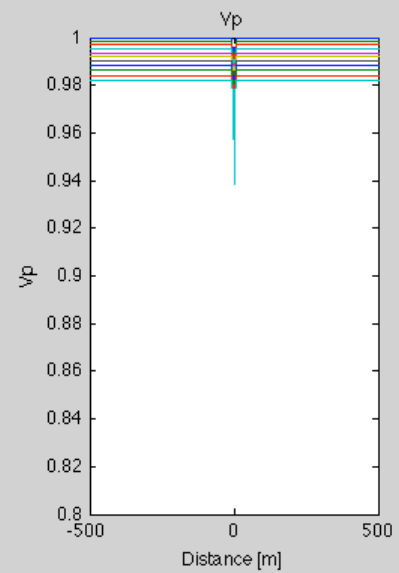
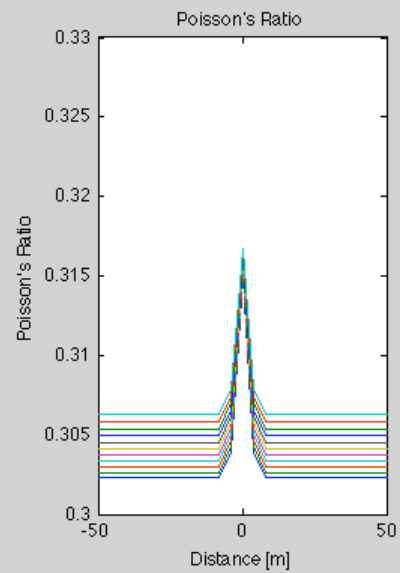
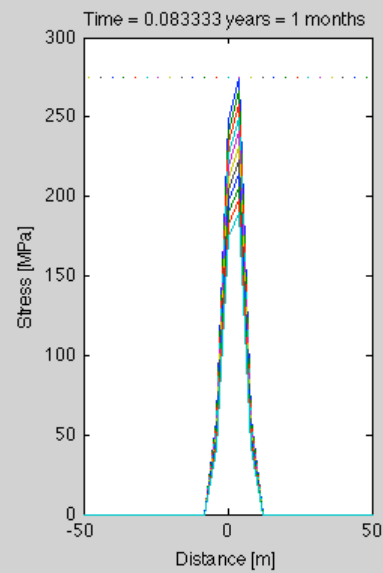
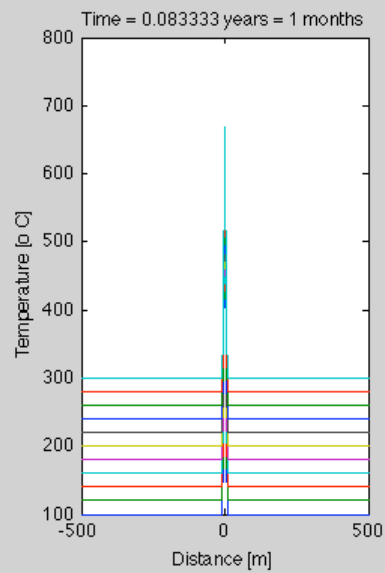
Note W, shift, S
propagation of
seismicity

A lot more to do!



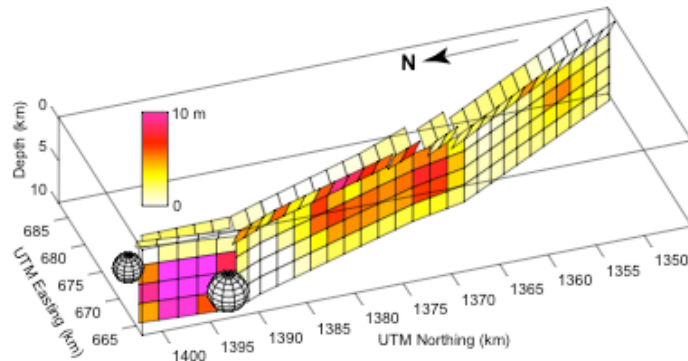


Kendall et al., in press



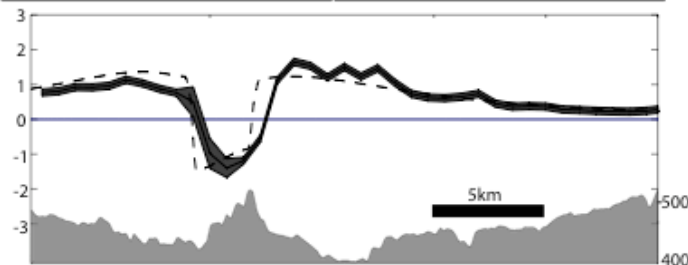
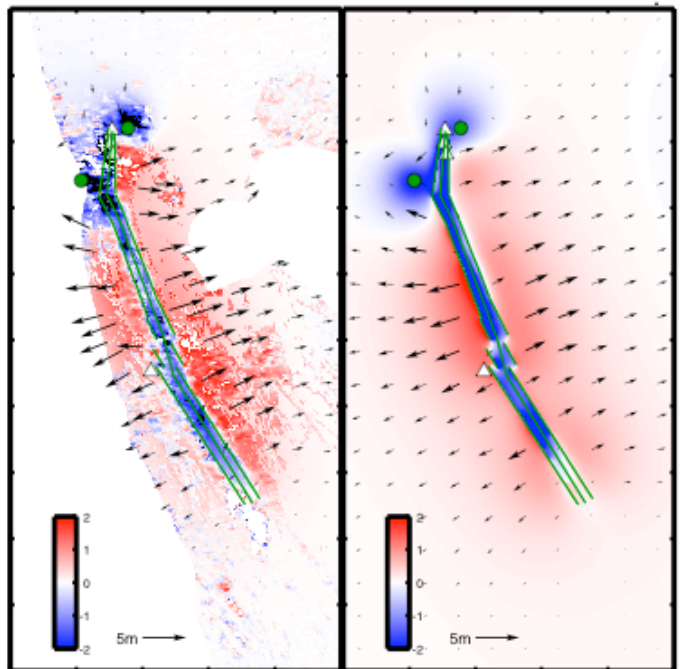
Infrastructure and capacity building –
sustainable solutions may come from:

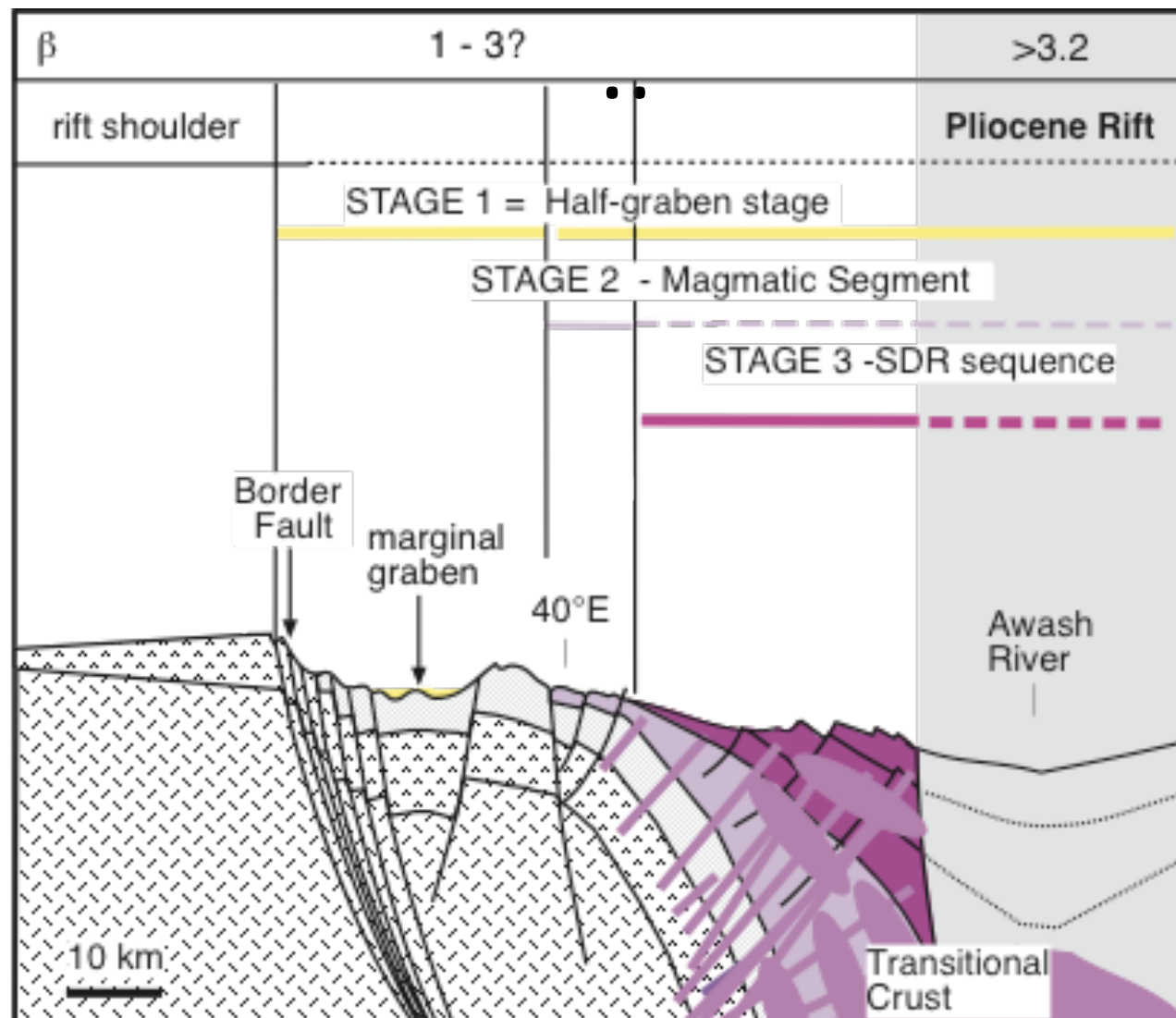
- Data archiving – resource and storage ?
- Software updates – sharing ‘clones’ ?
- Regular software training via internet/
training schools ?
- Visiting scholars ?
- Observatories as regional resources?
- Sandwich courses for PhD and MSc
research?



Wright et al., Nature, 2006

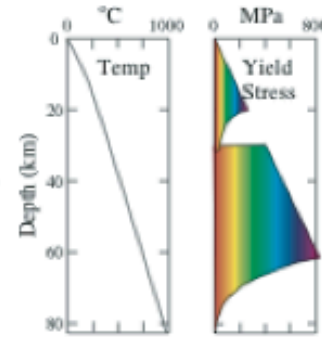
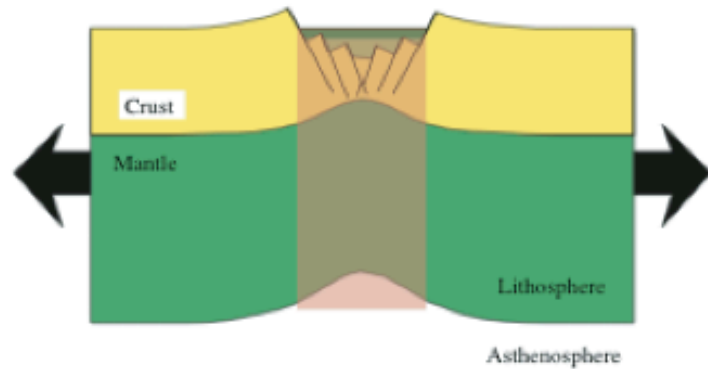
Model of interferometry
 2 Mogi sources
 ~ 8 m opening, 60 km-long
 dyke, <30% sourced from
 shallow magma chambers
 beneath Dabbahu and
 Gabho



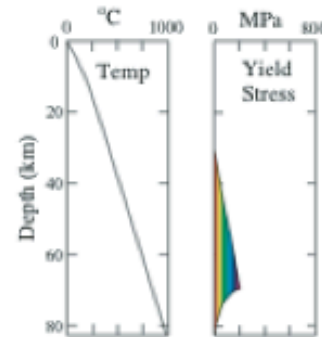
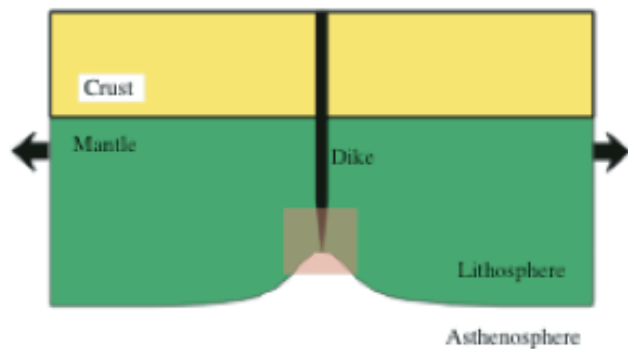


Wolfenden, 2003

Tectonic Stretching



Magmatic Extension



Mechanical stretching - differences relate to rheology assumed.

Add the additional buoyancy force of magma intrusion and associated heating, and breakup can occur at 1/10 the force required for Class A.

After Buck, 2004

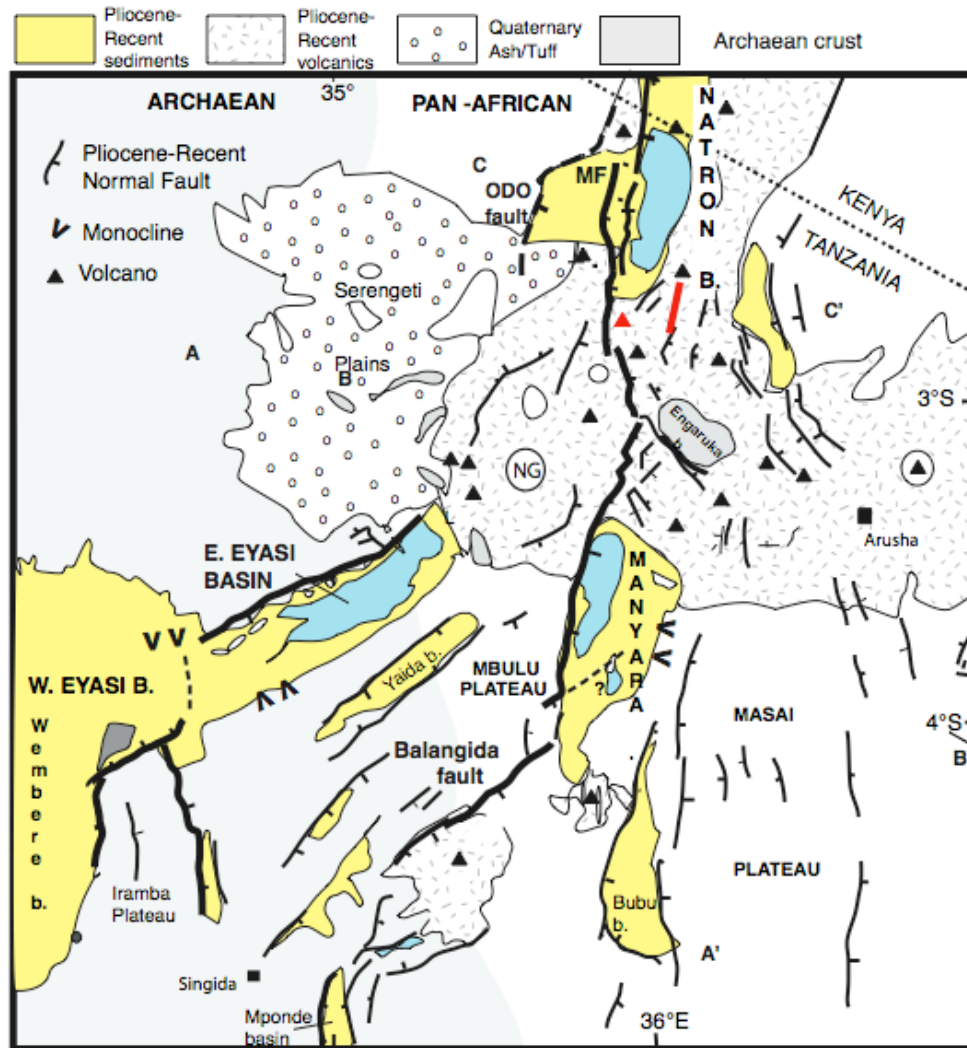
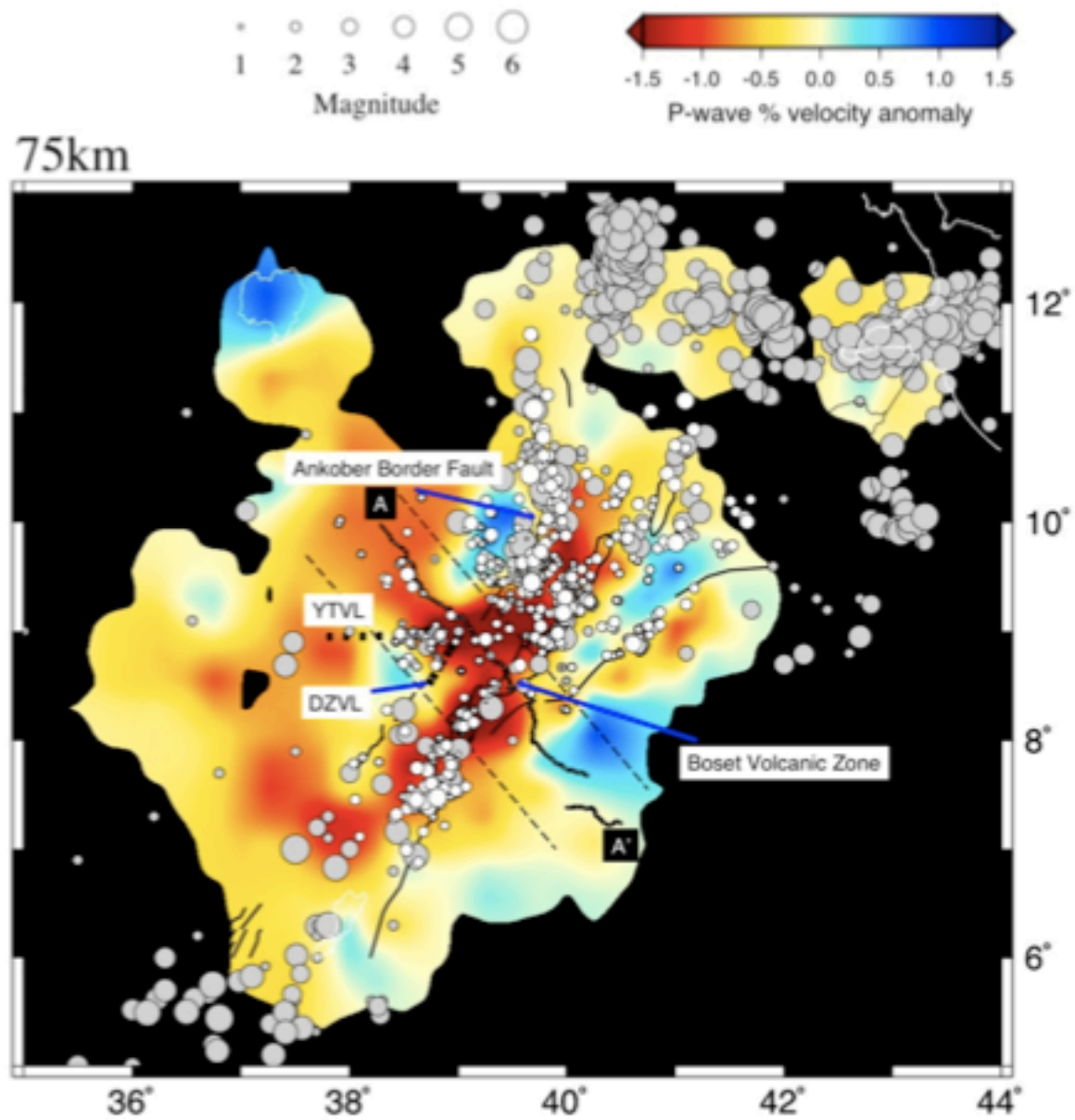


Figure 4. Ebinger et al.

Figure 2; Keir et al.



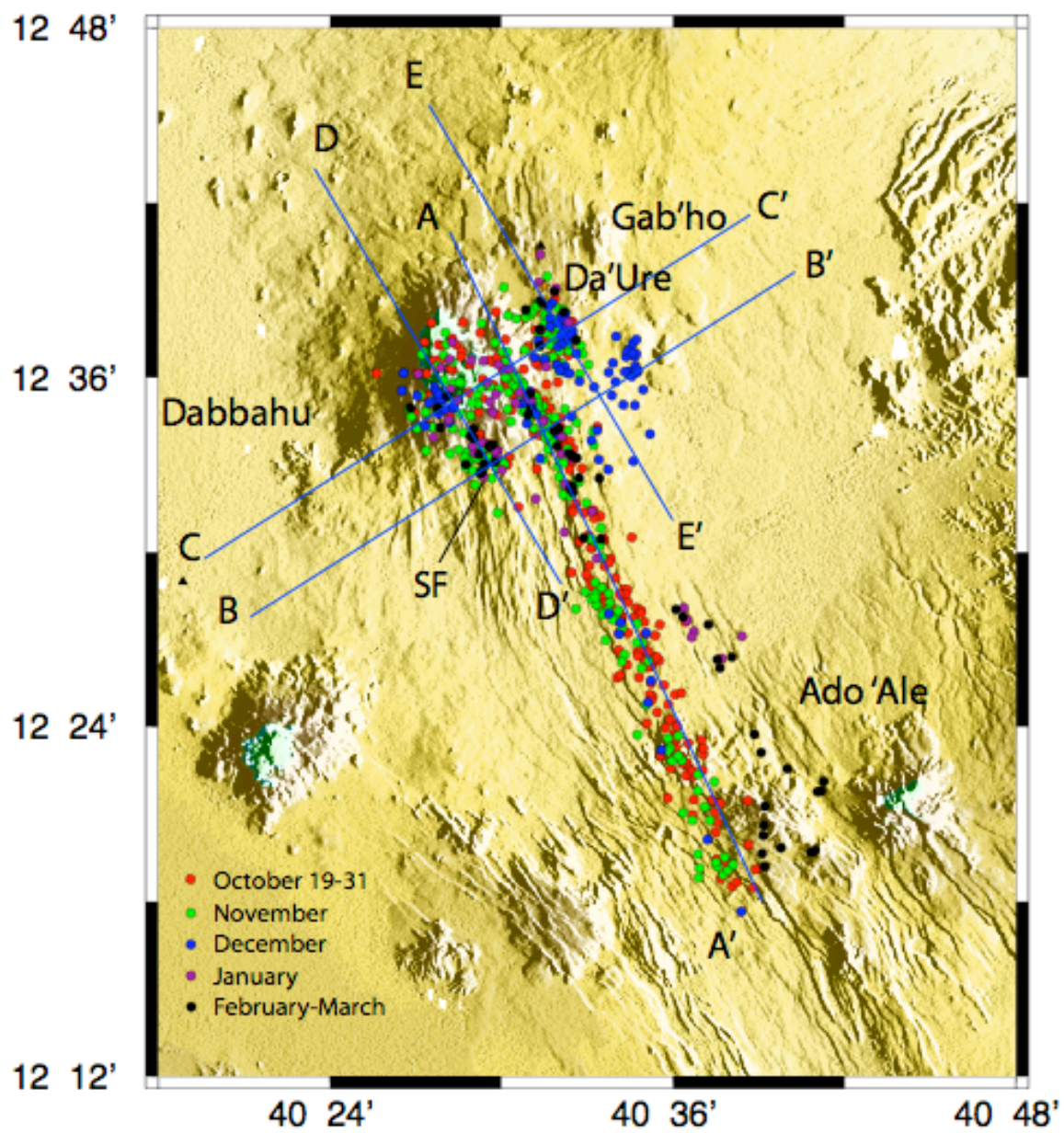
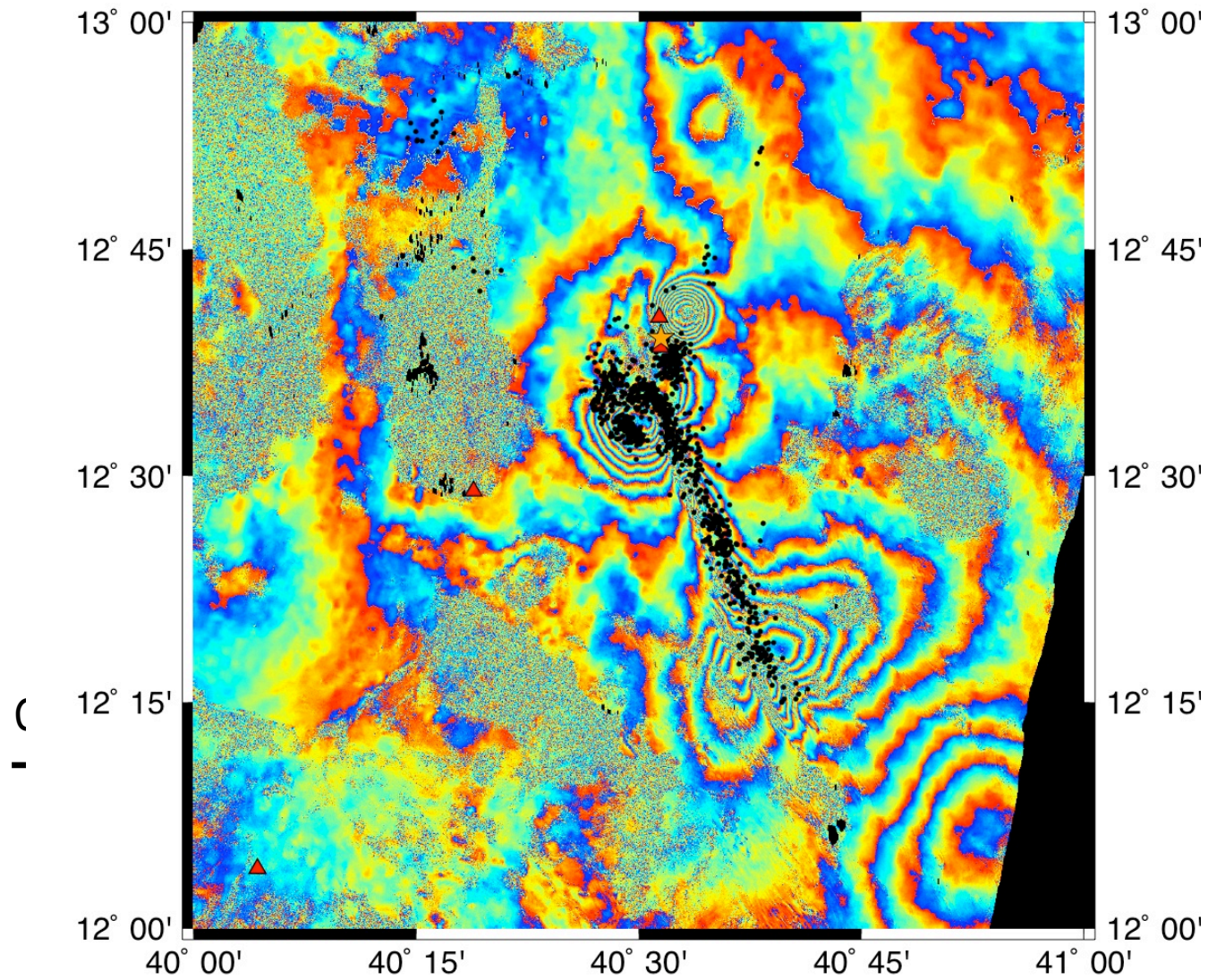


Fig.3, Ebinger et al.



Seismicity 18/10-30/4/06

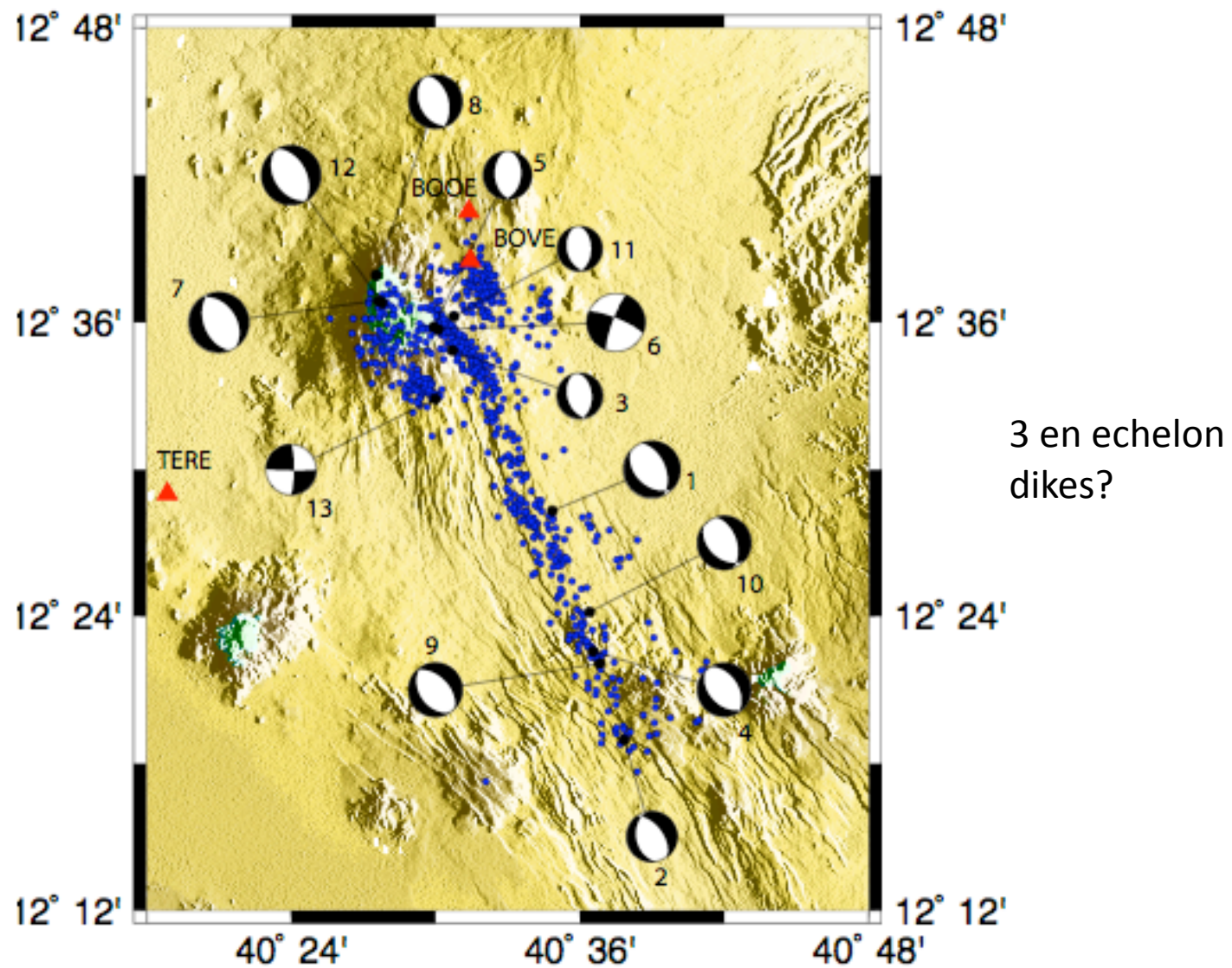
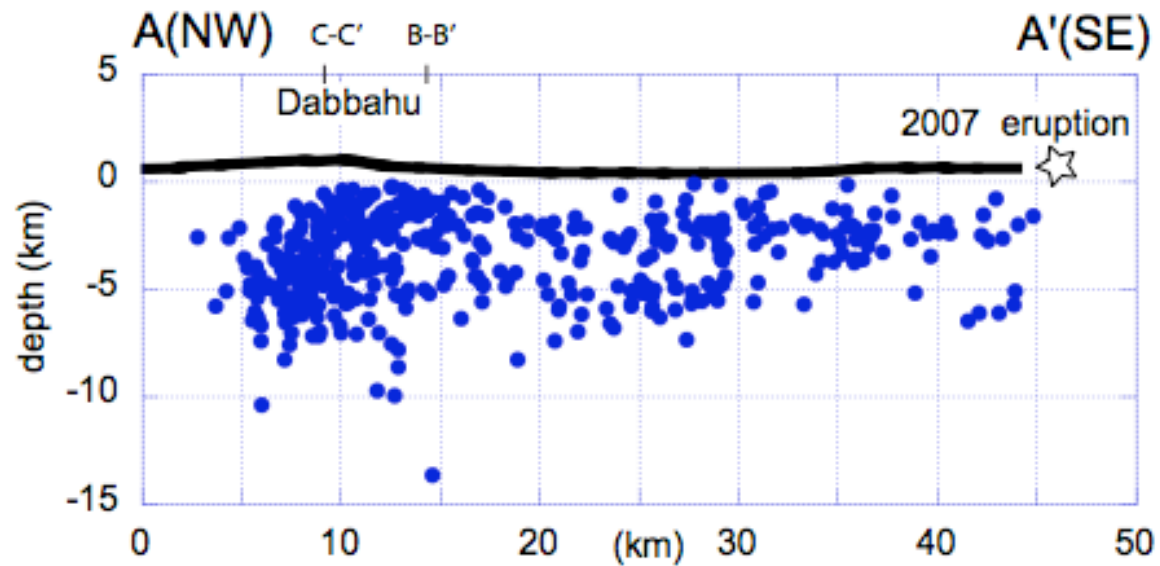


Fig.5, Ebinger et al.



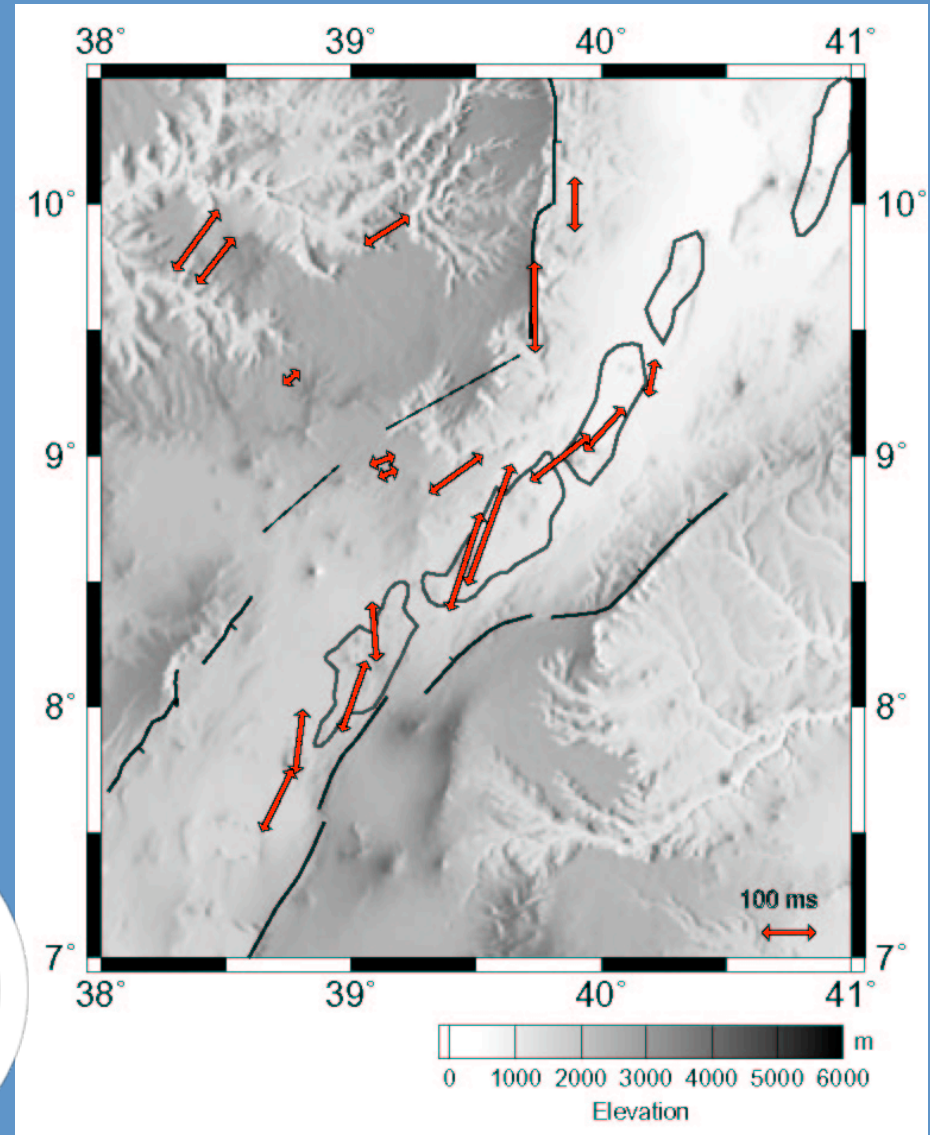
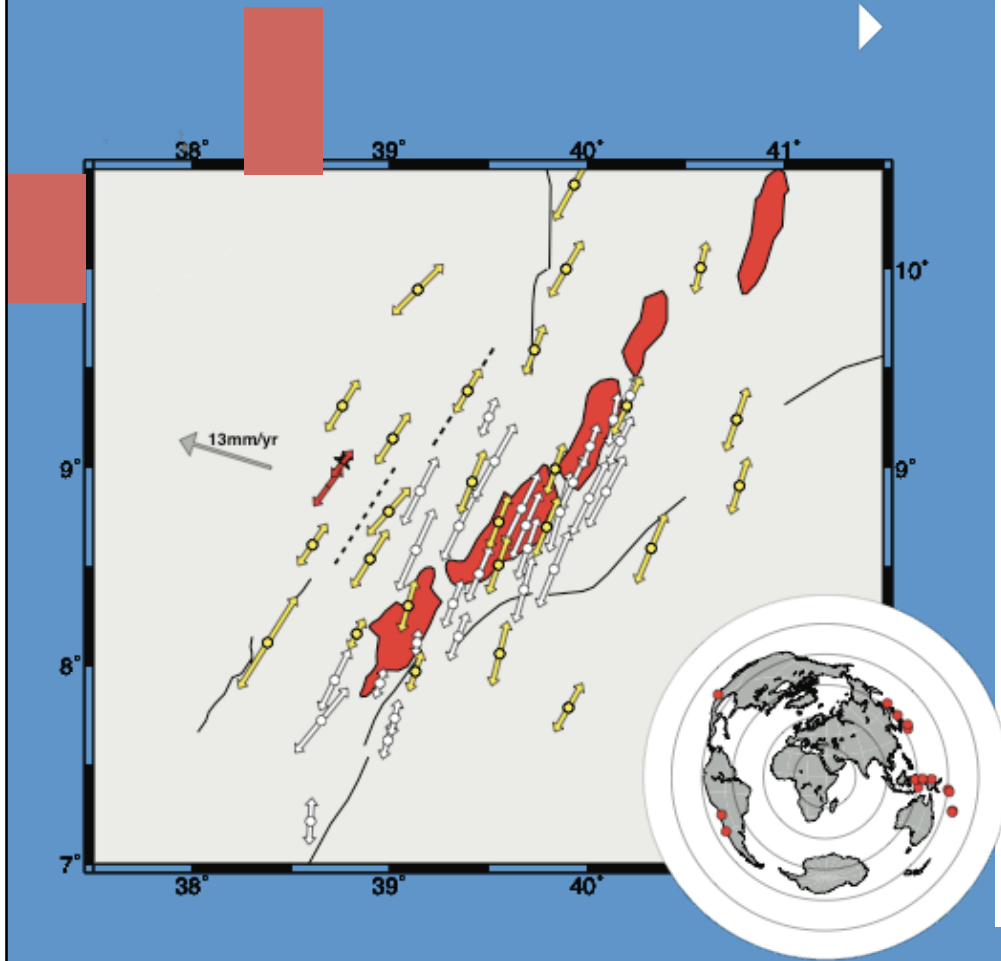
Post-intrusion seismicity
Thermal stress along dike walls; continued
dike intrusion?

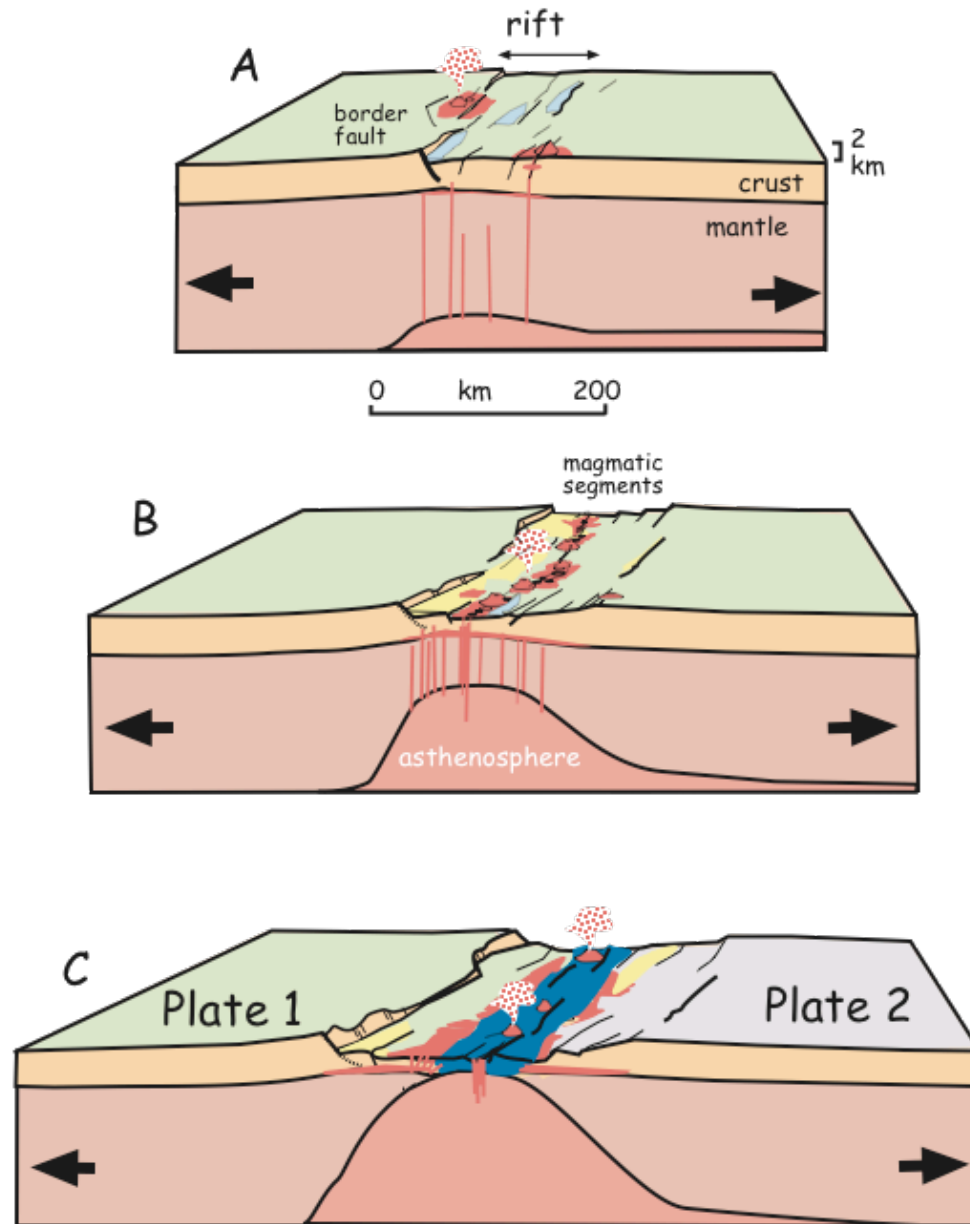
Figure 7, Ebinger et al.

Shear-wave splitting indicates melt-filled cracks that penetrate plate

Kendall et al.,
Nature 05 - > 2 s
SKS splitting

Keir et al.
GRL '05
4-6%
anisotropy





Rifting with Magma

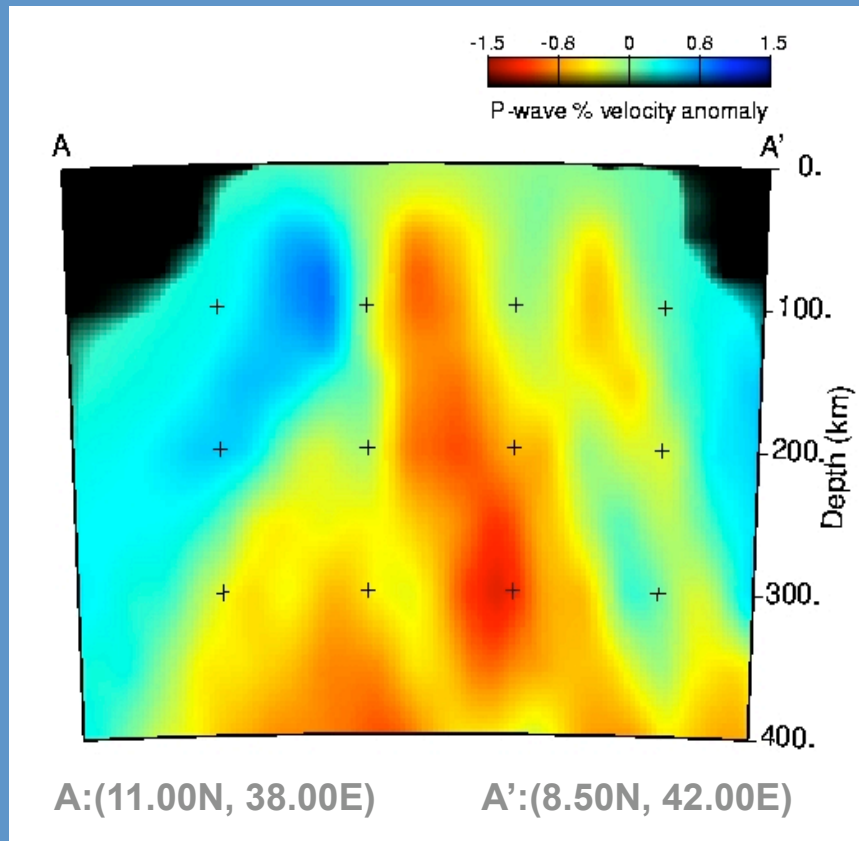
Stage 1 - Thick lithosphere, long repeat times; large stored stresses.

Stage 2: Strain localization to zones of magma intrusion rather than older border faults

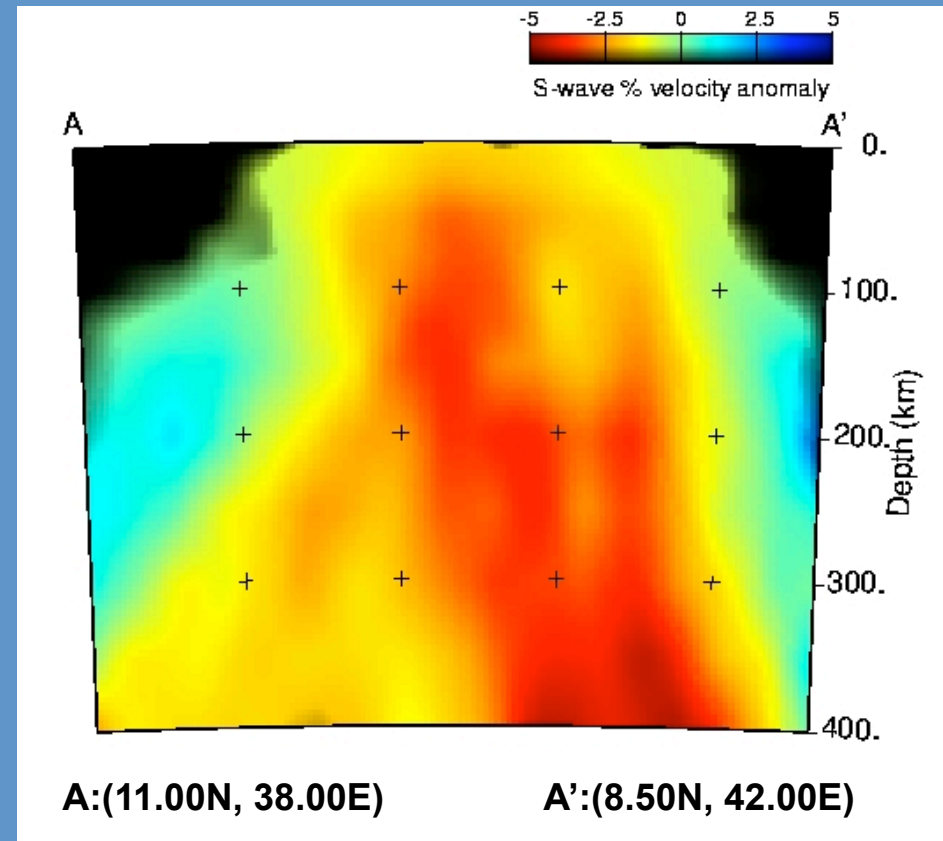
Stage 3: Onset of seafloor spreading similar to ultra-slow ridges - ridge jumps, 'off-axis' volcanism, and long repeat times.

Fig 3.

P



S



Ian Bastow, Graham Stuart, & Mike Kendall, U of Leeds

Bastow et al., GJI, 2005 – see also Benoit et al., 2006, Bastow et al., 2009

P

S