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A. Soloviev International Institute of Earthquake Prediction Theory and Mathematical Geophysics Moscow RUSSIAN FEDERATION

www.mitp.ru

Strada Costiera 11, 34151 Trieste, Italy - Tel.+39 040 2240 111; Fax +39 040 224 163 - sci_info@ictp.it

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International Institute of Earthquake Prediction Theory and Mathematical Geophysics Russian Academy of Sciences 84/32 Profsouznaya st., Moscow 117997 Russian Federation www.mitp.ru

I. INTRODUCTION

The problem under consideration is to determine in the region the areas where strong (with magnitude $M \ge M_0$ where M_0 is a threshold specified) earthquakes are possible. The detailed description of this problem, approaches to its solving, and a review of the results obtained for several regions are given by *Gorshkov et al.* (2003). The basic assumption is that strong earthquakes associate with morphostructural nodes, specific structures that are formed about intersections of fault zones. This gives possibility to apply the pattern recognition approach.

The nodes are considered as objects of recognition. They are identified by means of the morphostructural zoning and described by functions determined on the basis of the topographical, geological, geomorphological and geophysical data. When these functions are measured, the objects are represented by vectors with components, which are values of the functions.

The problem as the pattern recognition one is to divide the vectors into two classes: vectors D (Dangerous) and vectors N, which represent correspondingly the nodes where earthquakes with $M \ge M_0$ may occur and the nodes where only earthquakes with $M < M_0$ may occur. Application of the pattern recognition algorithms requires a training set of vectors, for which we know *a priori* the class they belong to. The training set is formed on the basis of the data on seismicity observed in the region. It consists of vectors D_0 and N_0 representing correspondingly the nodes where strong earthquakes occurred and the nodes, which are far from the known epicenters of such earthquakes.

II. FORMULATION OF THE PROBLEM AND THE MAIN STAGES OF ITS INVESTIGATION

Consider a selected magnitude cutoff M_0 that defines large earthquakes in the region under study. Roughly speaking, the problem of determining earthquake-prone areas aims at separating places of potential earthquakes into two parts, D where earthquakes with magnitude $M \ge M_0$ can happen and N where earthquakes with magnitude $M \ge M_0$ are impossible.

The first question arising in a strict formulation of the pattern recognition problem is how to select the region and the magnitude cutoff M_0 . The experience accumulated in *Gelfand et al.* (1972, 1973, 1974a, 1974b, 1976), *Zhidkov et al.* (1975), *Gvishiani et al.* (1978, 1987), *Caputo et al.* (1980), *Zhidkov and Kossobokov* (1980), *Gvishiani and Kossobokov* (1981), *Kossobokov* (1983), *Gvishiani and Soloviev* (1984), *Cisternas et al.* (1985), and *Gorshkov et al.* (1987) suggests the following heuristic criteria.

- The number of large earthquakes with $M \ge M_0$ in the region should be at least 10-20.
- The circles centered at epicenters of reported earthquakes with $M \ge M_0$ that have radii about the size of their source should not cover the entire region (otherwise, the problem has a trivial solution where the whole region is *D*).
- The region has to be tectonically uniform in sense of the similarity of possible causes of earthquakes with $M \ge M_0$.

These criteria establish certain limitations on the size of the region and the threshold M_0 . For instance, $M_0 = 5.0-6.0$ implies the linear size of a region of the order of hundreds kilometers, whereas for $M_0 = 7.0-7.5$ this size should be larger than a thousand kilometers. $M_0 = 8.0$ requires a region tens of thousands kilometers long. These limitations were met in practice, for example, in Italy, $M_0 = 6.0$ (*Caputo et al.*, 1980); in California, $M_0 = 6.5$ (*Gelfand et al.*, 1976); in South America and Kamchatka, $M_0 = 7.75$ (*Gvishiani and Soloviev*, 1984), and in the whole Circumpacific, $M_0 = 8.0$ (*Gvishiani et al.*, 1978). The experience accumulated in a decade confirmed that pattern recognition methods might reliably

distinguish earthquake-prone areas on different scales of lithospheric block hierarchy and in different seismic and tectonic environments (*Gelfand et al.*, 1972, 1973, 1974a, 1974b, 1976; *Zhidkov et al.*, 1975; *Gvishiani et al.*, 1978, 1987; *Caputo et al.*, 1980; *Zhidkov and Kossobokov*, 1980; *Gvishiani and Kossobokov*, 1981; *Kossobokov*, 1983; *Gvishiani and Soloviev*, 1984; *Cisternas et al.*, 1985; *Gorshkov et al.*, 1987).

When selecting the region and threshold magnitude M_0 , it is necessary to define the objects of recognition.

Gelfand et al. (1972) were the first who applied pattern recognition methods to determine earthquake-prone areas in the Pamirs and Tien Shan. Since then, several important improvements in such a determination have been developed, including a broader choice of natural objects for recognition. In general, one may consider three types of objects in a study of earthquake-prone areas: planar areas, segments of linear structures, and points.

Gelfand et al. (1972) used planar morphostructural nodes of the Pamirs and Tien Shan as candidates for earthquake-prone areas. At that time, even a formal definition of this structure that permits reproducible identification did not exist and was subject of further analysis by gemorphologists and mathematicians (*Alekseevskaya et al.*, 1977). However, because most fractional areas are characterized by multidirectional intensive tectonic movements, nodes essentially attract epicenters of large earthquakes. The fact that most earthquakes with $M \ge M_0$ in a region originate within nodes is a necessary precondition for using them as objects of recognition. *Ranzman* (1979) formulated the geomorphological basis that favors this precondition. *Gvishiani and Soloviev* (1981) suggested a statistical method for testing it in practice, even when the boundaries of nodes are not defined precisely.

In planar nodes, pattern recognition algorithms classify morphostructural node in the region either as a *D* node, which is prone to earthquakes with $M \ge M_0$, or as a *N* node, where strong earthquakes are not possible. Such a classification determines the area *D* as the union of all *D* nodes and the area *N* as the union of all *N* nodes. The remaining territories of the region complementary to the nodes are not assumed to be dangerous (they are rejected with a certain level of confidence by preconditioning strong earthquake – node association).

This natural choice of objects entails a difficult problem outlining the boundaries of morphostructural nodes. When the difficulty is overwhelming, one may try substituting the nodes with intersections of morphostructural lineaments as done by Gelfand et al. (1974b). Tracing lineaments and their intersections is much easier task for a geomorphologist that essentially delivers similar (though less complete) information on the most fractured places of multidirectional intensive tectonic movements. That is why intersections of morphostructural lineaments were commonly used for determining of earthquake-prone areas (Gelfand et al., 1974b, 1976; Zhidkov et al., 1975; Caputo et al., 1980; Zhidkov and Kossobokov, 1980; Gvishiani and Soloviev, 1984; Cisternas et al., 1985; Gorshkov et al., 1987; Gvishiani et al., 1987). The necessary precondition of using nodes as recognition objects is transformed to a hypothesis that epicenters of strong earthquakes originate near intersections of morphostructural lineaments (Gelfand et al., 1974b). This hypothesis is likely to be confirmed in a region if the following two conditions are valid: (1) the distance from all accurately determined epicenters of earthquakes with $M \ge M_0$ to the nearest intersection does not exceed a predefined distance r; (2) the area covered by circles of radius r centered in all intersections is a small part of the total area of the region. A statistical justification of the hypothesis can be obtained by using the algorithm developed by Gvishiani and Soloviev (1981).

Pattern recognition algorithms assign the vectors that describe intersections of lineaments to two classes: class D of intersections having vicinities prone to earthquakes with $M \ge M_0$ (D intersections) and class N. The classification of vectors determines the preimage of area D as the union of all vicinities of D intersections. The area N is the complement of area D in the union all vicinities of intersections. It is assumed that the remaining territories of the region complementary to all vicinities of intersections are not dangerous.

Usually, earthquakes are associated with segments of faults that they rapture. Therefore linear objects of recognition, like segments of active faults or fault zones, may seem most natural to many seismologists (*Gelfand et al.* (1976) give an excellent demonstration of how the problem is viewed differently). Pattern recognition algorithms divide linked linear objects into two classes: *D* segments capable of originating earthquakes with $M \ge M_0$ and *N* segments that are not.

Segments of linear structures were used as objects for recognition of earthquake-prone areas in California (*Gelfand et al.*, 1976), where the basic linear structure was San-Andreas fault, in the whole linear structure of Circumpacific seismic belt (*Gvishiani et al.*, 1978), and in the Western Alps (*Cisternas et al.*, 1985), where the segments of linear structures, forming a neotectonic scheme of the region were considered.

The usage of pattern recognition algorithms with learning necessitates an a priori selection of the training set W_0 , which is the union of two subjects that do not overlap: the training set D_0 from class D and the training set N_0 from class N. Such a selection of $W_0 = D_0 \cup N_0$ depends on the types of the objects for recognition. In the case of planar objects, all of those, including known epicenters of earthquake with $M \ge M_0$, form D_0 , whereas the subset N_0 consists of all remaining objects from W, $N_0 = W \setminus D_0$, or those of such objects that do not contain known epicenters of earthquakes with $M \ge M_0 - \delta$ (where $\delta > 0$ is usually 0.5 or about this value). It is necessary to emphasize that N_0 is not "pure" training set in the sense that some of its members belong to class D. In the first case, where $N_0 = W \setminus D_0$, the problem consists of distinguishing samples that spoil the purity of N_0 . Such a fussy type of learning highlights a specific difficulty in locating possible earthquake-prone areas by pattern recognition techniques.

It is natural to require the condition $D_0 \subseteq D$, where *D* denotes the vectors classified as belonging to class *D*. In other words, all places of strong earthquakes that are known should be recognized. When D_0 many vectors a part of it can be excluded from the training set and reserved to verify the reliability of the decision rule obtained.

When recognition objects are points, the training set D_0 is assembled from those that are situated at a distance not exceeded a certain fixed value r from the reported epicenters of earthquakes with $M \ge M_0$. The choice of r must satisfy the condition that the distance from most (practically all) of the well located epicenters of strong earthquakes in the region to the nearest recognition point is less than r. Naturally r scales with M_0 . For instance, *Zhidkov and Kossobokov* (1980) used r = 40 km for $M_0 = 6.5$ in the eastern part of Central Asia; *Gvishiani* and Soloviev (1984) chose r = 100 km for $M_0 = 7.75$ on the Pacific coast of South America. The training set N_0 consists of either all remaining points or those of them that are at a distance r_1 ($r_1 \ge r$) or longer from the epicenters of earthquakes with $M \ge M_0 - \delta$ ($\delta > 0$). In this case the training set N_0 also can contain points that are potentially from class D.

There is a certain difficulty when recognition objects are points; one epicenter can be attributed to several objects if its distance to each of them is *r* or less. In such case the training set D_0 may have some objects from class *N*. Algorithm CLUSTERS, which takes into account this specific feature of the training set D_0 is used to overcome this difficulty. In case of ambiguity, the condition that $D_0 \subseteq D$ is changed by another natural one: each epicenter of an earthquake with $M \ge M_0$ has a point *D* object at a distance *r* or less.

When recognition objects are linear segments, the training set D_0 assembles those containing a projection of an epicenter of a strong earthquake. The training set N_0 is either N_0 = $W \setminus D_0$ or contains segments from W that are not neighbors of D_0 . Another way to form N_0 is to exclude those segments from $W \setminus D_0$ that contain a projection of an epicenter of an earthquake with $M \ge M_0 - \delta$ (where $\delta > 0$ is a parameter). As a rule, there is a unique projection of an epicenter that does not create ambiguity in selecting D_0 : therefore, it is natural to require that $D_0 \subseteq D$. Pattern recognition algorithms operate with vectors of functions representing natural recognition objects. As far as the earthquake-prone areas are considered, it appears natural to use the functions describing, either directly or indirectly, the intensity of recent tectonic activity at the locality of each object. The accumulated experience in recognizing earthquake-prone areas has established the following functions as typical:

- a multitude of functions describing topography: maximum (H_{max}) and minimum (H_{min}) altitudes above sea level inside the object area, altitude range ΔH ; dominating combination of geomorphological structures in the object's vicinities, percentage of the object's area with existing Paleogene Quaternary sediments, etc.;
- functions describing the complexity of geomorphological and neotectonic network of structures: number of lineaments forming the object, the highest rank of lineament among those which form the object, etc.;
- functions describing gravitational field anomalies.

In case of planar objects the sense of "area" is obvious. When objects are points the area is a circle of the same radius for all objects centered at an object. When objects are linear segments the area is a circle of the same radius for all objects centered at the middle of a segment. Planar objects may have various areas and the area of an object may be used as one of functions.

In principle, all available information related directly or indirectly to the level of seismic activity can be used to characterize objects. The only necessary precondition for a function is availability of uniform measurements across the entire region under consideration. After measuring selected functions for all the objects, they are converted to vectors $\mathbf{w}^i = \{w_1^i, w_2^i, ..., w_m^i, \}, i = 1, 2, ..., n$, where *m* is the total number of functions, *n* is the total number of objects in *W*, and w_k^i is the value of the *k*-th function measured for the *i*-th object.

The pattern recognition algorithms, which are used to investigate the problem, work in a binary vectors space. Their application requires a transformation of vectors that describe natural recognition objects into binary ones.

Given the training set of vectors $W_0 = D_0 \cup N_0$, a pattern recognition algorithm determinates a classification $W = D \cup N$ where D and N are sets of vectors of classes D and N, respectively. As pointed above, the resulting classification should satisfy certain conditions, like $D_0 \subseteq D$ for planar objects. To avoid a trivial solution when all places considered belong to D, the following condition is usually introduced:

 $|D| \leq \beta |W|,$

where |D| and |W| stand for the numbers of objects in sets *D* and *W*, respectively; and β , $0 < \beta < 1$, is a real constant, which sets an a priori upper bound for the fraction of *D* vectors in *W*. The value and justification of β must result from an expert evaluation of geological, seismological, and other available information on the region.

The quality and reliability of a classification can be verified by control tests. If successful, such test favors the classification that actually divides the region into earthquakeprone areas and areas where earthquakes with $M \ge M_0$ are not likely. Usually, pattern recognition of earthquake-prone areas involves a small sample of natural objects whose size does not allow reserving a control set for verification. Nevertheless, certain verification of the classification can be achieved by the comprehensive analysis of the result and additional information that was not used initially, of which the most important are data on epicenters of large earthquakes, e.g., noninstrumental, either historical or paleoseismological.

Classifications that are not satisfactory and have no meaningful interpretation are usually not reported. To get a satisfactory classification, a researcher can perform several cycles of trial and error through the following stages of recognition:

• definition of the region under study and the magnitude cutoff attributed to earthquakeprone areas;

- choice of the natural recognition objects;
- selection of the training set $W_0 = D_0 \cup N_0$;
- description of objects as vectors;
- discretization and coding of the functions;
- classification of vector space $W = D \cup N$ by a pattern recognition algorithm;
- evaluation of the reliability of classification from control tests;
- interpretation of the classification $W = D \cup N$ as a division of the region into earthquakeprone and other areas;
- generalization of geological and geomorphological interpretation of classification and the rules used to obtain it.

After the definition of D and N areas in the region territory it is advisable to do a statistical analysis of the locations of the known epicenters of earthquakes with $M < M_0$ relative to the located areas (as, e.g., in *Kossobokov and Soloviev*, 1983). The result of such comparison can lead, in principle, to the conclusion that the obtained D and N areas are actually earthquake-prone areas for earthquakes with $M \ge M'_0$ where M'_0 is a smaller magnitude threshold than M_0 .

III. RECOGNITION OF EARTHQUAKE-PRONE AREAS AROUND THE ADRIA MARGIN IN PENINSULAR ITALY AND SICILY

The problem of recognition of places around the Adria margin in peninsular Italy and Sicily where earthquakes with $M \ge 6.0$ may occur (*Gorshkov et al.*, 2002) is briefly considered below.

The intersections of the morphostructural lineaments obtained as the result of the morphostructural zoning of the region under consideration are objects of pattern recognition. The scheme of the morphostructural zoning of the region and the objects are shown in Fig. 7. The total number of objects in the set *W* is 146. The problem is to classify these objects into two classes: objects where earthquakes with $M \ge 6.0$ may occur (class *D*) and objects where earthquakes with $M \ge 6.0$ are impossible (class *N*).

Two earthquake catalogues NT 4.1.1. (*Camassi and Stucchi*, 1997) and CCI-1996 (*Peresan et al.*, 1997), covering the entire region and containing events from 1000 to 1997 have been used to select the $M \ge 6.0$ quakes recorded in the region. Although these catalogues sometimes exhibit different values for the same parameters (chiefly magnitude) for the same events, they are the most complete sources on the seismic history of Italy. The training set D_0 includes intersections, hosting earthquakes with $M \ge 6.0$ in both catalogues. The epicenters of these earthquakes are shown in Fig. 8 by dots. The intersections hosting earthquakes with $M \ge 6.0$, in at least one of the two catalogues used, and the intersections situated in flat areas of low seismicity (Adriatic foreland and Tyrrhenian shelf) are not included both in D_0 and N_0 training sets. The remaining intersections are assigned to the set N_0 . As a result, 24 intersections (11, 26, 27, 43, 45, 51, 59, 61, 70, 73, 74, 84, 85, 90, 92, 95, 109, 112, 117, 118, 123, 128, 129, and 144) constitute D_0 and 66 intersections (1, 2, 4, 5, 6, 8, 10, 12, 13, 19, 29, 30, 33, 34, 35, 36, 37, 40, 42, 44, 46, 47, 49, 53, 56, 57, 60, 62, 72, 75, 76, 77, 81, 83, 89, 91, 94, 96, 98, 99, 101, 102, 104, 105, 107, 108, 110, 111, 114, 115, 116, 124, 125, 126, 131, 132, 133, 134, 135, 137, 138, 139, 140, 143, 145, 146) constitute N_0 .

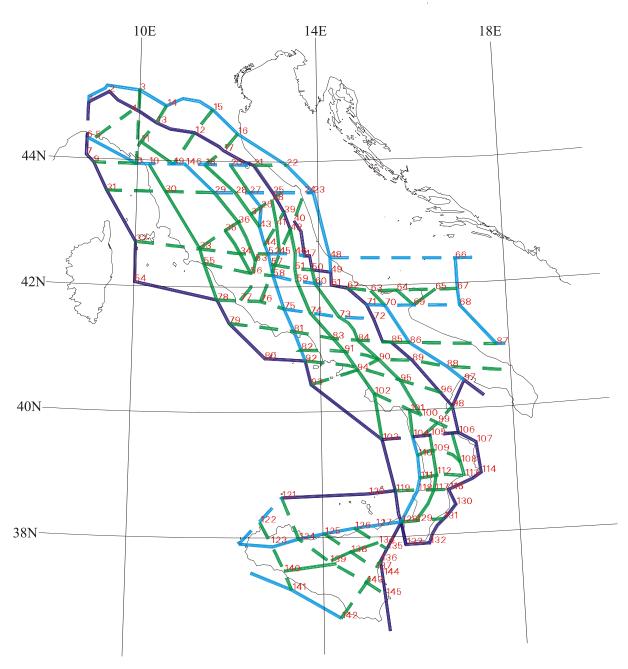


FIGURE 7 Morphostructural map around the Adria margin in peninsular Italy and Sicily. Continuous lines are the longitudinal lineaments, discontinuous ones are the transverse lineaments. Intersections of lineaments (objects of recognition) are numbered from 1 to 146.

Table 1 lists of functions, which describe the objects. The components of vectors \mathbf{w}^i are the values of these functions. The values of the functions have been measured from topographic, geological, gravity and morphostructural maps within the areas of radius of 25 km around the intersection of the lineaments. The discretization thresholds for the functions are also given in Table 1. Except for the morphological function (Mor), their binary coding is *S* type, for Mor it is *I* type.

The value of β , which sets an a priori upper bound for the fraction of *D* vectors in *W*, has been estimated as 0.6. Therefore classifications with $|D| \le 0.6 |W|$ are considered only.

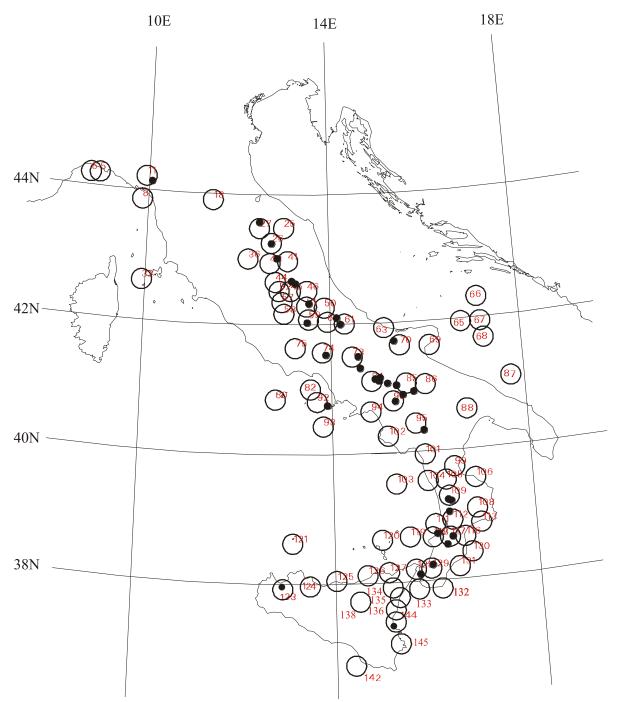


FIGURE 8 Result of the recognition of the intersections prone to earthquakes with $M \ge 6.0$. Dots are the epicenters of earthquakes with $M \ge 6.0$ in both catalogues used. Circles are the areas of radius of 25 km around the intersection recognized to be prone to earthquakes with $M \ge 6.0$. Numbering of intersections is the same as in Fig. 7.

After preliminary analysis six functions have been left in the binary codes of the objects: relief energy (Δ H), gradient of topography (Δ H/L), minimal value of Bouguer anomaly (Bmin), highest rank of lineament (HR), distance to the nearest second rank lineament (D2), and morphology (Mor). The result of the learning stage is shown in Table 2. The characteristic traits of *D* and *N* intersections have been obtained by "CORA-3" with the

following values of the thresholds: $k_1 = 4$, $\bar{k}_1 = 2$, $k_2 = 13$, $\bar{k}_2 = 0$. The traits are given in the table as conjunctions of inequalities in the values of the functions of the intersections.

Functions	Thresholds of discretization		
	first	second	
A)Topographic functions			
Maximum topographic altitude, <i>m</i> (Hmax)	1500		
Minimum topographic altitude, m (Hmin)	-230	80	
Relief energy, m (Δ H) (Hmax - Hmin)	1500	2000	
Distance between the points Hmax and Hmin, km (L)	35		
Slope, $(\Delta H/L)$	40	65	
B) Geological function			
The portion of soft (quaternary) sediments, %, (Q)	1	5	
C) Gravity functions			
Maximum value of Bouguer anomaly, <i>mGal</i> ,(Bmax)	10	47	
Minimum value of Bouguer anomaly, mGal, (Bmin)	- 46	7	
Difference between Bmax and Bmin, $mGal$, (ΔB)	44	66	
D) Functions from the morphostructural map			
The highest rank of lineament in an intersection,			
(HR)	1		
Number of lineaments forming an intersection, (NL)	2		
Distance to the nearest 1st rank lineament, km, (D1)	0	50	
Distance to the nearest 2nd rank lineament, km, (D2)	0	50	
Distance to the nearest intersection, km, (Dn)	23	30	
E) Morphological function (Mor)			
This parameter is equal to one of the following six			
values in accord with the morphology within vicinity	2	4	
of each intersection:			
1 - mountain and plain (m/p)			
2 - mountain and piedmont (m/pd)			
3 - mountain and mountain (m/m)			
4 - piedmont and plain (pd/p)			
5 - piedmont only (pd)			
6 - plain only (p)			

 TABLE 1 Functions describing the objects of recognition

TABLE 2 Characteristic traits selected b	y algorithm CORA-3
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#	$\Delta H, m$	$\Delta H/L$	Bmin, <i>mGal</i>	HR	D2, <i>km</i>	Mor
	D traits					
1	≤ 2000	> 65				m/m <i>or</i> pd/p
2	≤ 1500	>40				m/m <i>or</i> pd/p
3	≤ 2000		>7			m/m <i>or</i> pd/p
	N traits					
1		>40				<i>not</i> (m/m <i>or</i> pd/p)
2	≤ 1500	≤ 40	≤ 7			
3			≤ 7			<i>not</i> (m/m <i>or</i> pd/p)
4				≤ 1		<i>not</i> (m/m <i>or</i> pd/p)
5					> 50	<i>not</i> (m/m <i>or</i> pd/p)

The classification has been made with $\Delta = 0$, i.e. a node is assigned to the *D* set, if the difference between the number of *D*- and *N*-traits, which a given node possess, is greater or equal to 0. The classification of the intersections is shown in Fig. 8: 81 intersections (55% of the total number of intersections) are assigned to class *D*, and the remaining 65 to *N*. Class *D* includes all D_0 , 28 intersections from N_0 , and 29 intersections from outside the training sets.

IV. REVIEW OF THE RESULTS ON RECOGNITION OF EARTHQUAKE-PRONE AREAS

Table 3 contains a list of regions where the earthquake-prone areas have been determined.

TABLE 3 Regions where the earthquake-prone areas have been determined					
Region	M_0	Reference			
The Western Alps	5.0	Cisternas et al. (1985)			
The Pyrenees	5.0	Gvishiani et al. (1987)			
The Greater Caucasus, intersections	5.0	Gvishiani et al. (1988)			
The Greater Caucasus, nodes	5.5	Gorshkov et al. (2003)			
Italy	6.0	<i>Caputo et al.</i> (1980);			
		Gorshkov et al. (2002)			
The Alps and Dinarides	6.0	Gorshkov et al. (2004)			
Tien Shan and Pamirs	6.5	Gelfand et al. (1972)			
Balkans, Asia Minor, Transcaucasia	6.5	Gelfand et al. (1974a)			
California and Nevada	6.5	Gelfand et al. (1976)			
Himalayas	6.5	<i>Bhatia et al.</i> (1992)			
	7.0	Bhatia et al. (1994)			
Andes of South America	7.75	Gvishiani and Soloviev (1984)			
Circumpacific seismic belt	8.2	Gvishiani et al. (1978)			

TABLE 3 Regions where the earthquake-prone areas have been determined

Table 4 summarizes up to 2003 the comparison between the location of epicenters of strong earthquakes occurred in these regions after completing the recognition and the results of the earthquake-prone areas determination (*Gorshkov et al.*, 2003). One can see from this table that only 4 of 71 strong earthquakes have occurred in N-objects and 8 strong earthquakes have occurred outside the objects of recognition. Note that 18 strong earthquakes have occurred in D-objects that did not belong to the training set D_0 . Such D-objects are marked by *.

Region	M_0	Result obtained	Total number of	Occurred in		Out of recognition	
		in	strong	D (D*)-	N-	objects	
			earthquakes	objects	objects		
The Western Alps	5.0	1985	5	4 (1)	1	-	
The Pyrenees	5.0	1987	2	1	1	-	
The Greater	5.0	1988	14	11 (3)	1	2	
Caucasus							
Italy	6.0	1979	5	3 (1)	-	2	
Tien Shan and	6.5	1972	6	4 (1)	-	2	
Pamirs							

TABLE 4 Summary of the test of earthquake-prone areas determination

Balkans, Asia	6.5	1974	20	19 (5)	1	-
Minor,						
Transcaucasia						
California and	6.5	1976	15	13 (5)	-	2
Nevada						
Himalayas	6.5	1992	2	2 (1)	-	-
Andes of South	7.75	1982	2	2 (1)	-	-
America						
Total	Total			59 (18)	4	8

REFERENCES

- Alekseevskaya, M.A., A.M.Gabrielov, A.D.Gvishiani, I.M.Gelfand, and E.Ya.Ranzman (1977). Formal morphostructural zoning of mountain territories. *J. Geophys*, **43**, 227-233.
- Bhatia,S.C., T.R.K.Chetty, M.Filimonov, A.Gorshkov, E.Rantsman, and M.N.Rao (1992). Identification of potential areas for the occurrance of strong earthquakes in Himalayan arc region. *Proc. Indian Acad. Sci. (Earth Planet. Sci.)*, **101**(4): 369-385.
- Bhatia,S.G., M.N.Rao, T.R.K.Chetty, E.Ya.Rantsman, A.I.Gorshkov, M.B.Filimonov, and N.V.Shtock (1994). Recognition of earthquake-prone areas. XIX. The Himalaya, M ≥ 7.0. In V.I.Keilis-Borok and G.M.Molchan (eds), *Theoretical Problems of geodynamics and Seismology*. Moscow: 280-287 (Comput. Seismol.; Iss. 27, in Russian).
- Camassi,R., and M.Stucchi (1997). NT 4.1, Un catalogo parametrico di terremoti di area Italiana al di Sopra della soglia di danno, Open data file. Consiglio Nazionale delle Ricerche GNDT.
- Caputo, M., V.Keilis-Borok, E.Oficerova, E.Ranzman, I.Rotwain, and A.Solovjeff (1980). Pattern recognition of earthquake-prone areas in Italy. *Phys. Earth Planet Int.*, 21: 305-320.
- Cisternas, A., P.Godefroy, A.Gvishiani, A.I.Gorshkov, V.Kossobokov, M.Lambert, E.Ranzman, J.Sallantin, H.Soldano, A.Soloviev, and C.Weber (1985). A dual approach to recognition of earthquake prone areas in the western Alps. *Annales Geophysicae*, **3**, 2: 249-270.
- Gelfand,I.M., Sh.Guberman, M.L.Izvekova, V.I.Keilis-Borok, and E.Ia.Ranzman (1972). Criteria of high seismicity determined by pattern recognition. In A.R.Ritsema (ed.), *The Upper Mantle. Tectonophysics*, **13** (1-4): 415-422.
- Gelfand,I.M., Sh.I.Guberman, M.P.Zhidkov, M.S.Kaletskaya, V.I.Keilis-Borok, and E.Ja.Ranzman (1973). Transfer of high seismicity criteria from the East of Central Asia to Anatolia and adjacent regions. *Doklady Academii Nauk SSSR*, 210, 2: 327-330 (in Russian).
- Gelfand,I.M., Sh.I.Guberman, M.P.Zhidkov, M.S.Kaletskaya, V.I.Keilis-Borok, E.Ja.Ranzman, and I.M.Rotwain (1974a). Identification of sites of possible strong earthquake occurrence. II. Four regions of Asia Minor and South Eastern Europe. In V.I.Keilis-Borok (ed.), *Computer Analysis of Digital Seismic Data*. Moscow, Nauka: 3-40 (Comput. Seismol.; Iss. 7, in Russian).
- Gelfand,I.M., Sh.I.Guberman, M.P.Zhidkov, V.I.Keilis-Borok, E.Ja.Ranzman, and I.M.Rotwain (1974b). Identification of sites of possible strong earthquake occurrence.
 III. The case when the boundaries of disjunctive knots are unknown. In V.I.Keilis-Borok (ed.), *Computer Analysis of Digital Seismic Data*. Moscow, Nauka: 41-65 (Comput. Seismol.; Iss. 7, in Russian).
- Gelfand,I.M., Sh.A.Guberman, V.I.Keilis-Borok, L.Knopoff, F.Press, I.Ya.Ranzman, I.M.Rotwain, and A.M.Sadovsky (1976). Pattern recognition applied to earthquake epicenters in California. *Phys. Earth Planet. Inter.*, **11**: 227-283.
- Gorshkov,A.I. G.A.Niauri, E.Ya.Ranzman and A.M.Sadovsky (1987). Use of gravimetric data for recognition of places of possible occurance of strong earthquakes in the Great Caucasus. In V. I. Keilis-Borok and A. L. Levshin (eds), *Theory and Analysis of Seismological Information*, Allerton Press Inc, New York: 117-123 (Comput. Seismol., volume 18).
- Gorshkov, A.I., G.F.Panza, A.A.Soloviev, and A.Aoudia (2002). Morphostructural zonation and preliminary recognition of seismogenic nodes around the Adria margin in peninsular Italy and Sicily. J. Seismol. & Earthquake Engineering, **3**, 2: 1-24.

- Gorshkov, A., V.Kossobokov, and A.Soloviev (2003). Recognition of Earthquake-Prone Areas. In V.I.Keilis-Borok and A.A.Soloviev (eds), *Nonlinear Dynamics of the Lithosphere and Earthquake Prediction*. Springer-Verlag, Berlin-Heidelberg: 239-310.
- Gorshkov,A.I., G.F.Panza, A.A.Soloviev, and A.Aoudia (2004). Identification of seismogenic nodes in the Alps and Dinarides. *Bolletino della Societa Geologica Italiana*, **123**(1): 3-18.
- Gvishiani, A.D., A.V.Zelevinsky, V.I.Keilis-Borok, and V.G.Kossobokov (1978). Study of the violent earthquake occurrences in the Pacific Ocean Belt with the help of recognition algorithms. *Izvestiya Acad. Sci. USSR. Physics of the Earth*, 9: 31-42 (in Russian).
- Gvishiani, A.D. and V.G.Kossobokov (1981). On foundations of the pattern recognition results applied to earthquake-prone areas. *Izvestiya Acad. Sci. USSR. Physics of the Earth*, 2: 21-36 (in Russian).
- Gvishiani,A., and A.Soloviev (1981). Association of the epicenters of strong earthquakes with the intersections of morphostructural lineaments in South America. In V. I. Keilis-Borok and A. L. Levshuin (eds), *Interpretation of Seismic Data - Methods and Algorithms*, Allerton Press Inc., New York: 42+ (Comput. Seismol., volume 13).
- Gvishiani,A.D., and A.A.Soloviev (1984). Recognition of places on the Pacific coast of the South America where strong earthquakes may occur. *Earthq. Predict. Res.*, 2: 237-243.
- Gvishiani, A., A.Gorshkov, V.Kossobokov, A.Cisternas, H.Philip, and C.Weber (1987). Identification of seismically dangerous zones in the Pyrenees. *Annales Geophysicae*, 5B(6): 681-690.
- Gvishiani, A., A.Gorshkov, E.Rantsman, A.Cisternas, and A.Soloviev (1988). Identification of Earthquake-prone Areas in the Regions of Moderate Seismicity, Nauka, Moscow, (in Russian).
- Kossobokov, V.G. (1983). Recognition of the sites of strong earthquakes in East Central Asia and Anatolia by Hamming's method. In V.I.Keilis-Borok and A.L.Levshin (eds), *Mathematical Models of the Structure of the Earth and the Earthquake Prediction*, Allerton Press Inc, New York: 78-82 (Comput. Seismol., volume 14).
- Kossobokov,V.G., and A.A.Soloviev (1983). Disposition of epicenters of earthquakes with *M* ≥ 5.5 relative to the intersection of morphostructural lineaments in the East Central Asia. In V.I.Keilis-Borok and A.L.Levshin (eds), *Mathematical Models of the Structure of the Earth and the Earthquake Prediction*, Allerton Press Inc, New York: 75-77 (Comput. Seismol., volume 14).
- Peresan, A., G.Costa, and F.Vaccari (1997). CCI1996: the current catalogue of Italy. *Internal Report* IC/IR/97/9. Miramare-Trieste, April 1997, 422 p.
- Ranzman, E.Ia. (1979). *Places of Earthquakes and Morphostructures of Mountain Countries*. Nauka, Moscow (in Russian).
- Zhidkov,M.P., I.M.Rotwain, A.M.Sadovsky (1975). Recognition of places where strong earthquakes may occur. IV. High-seismic intersections of lineaments of the Armenian upland, the Balkans and the Aegean Sea basin. In V.I.Keilis-Borok (ed.), *Interpretation of Seismology and Neotectonics Data*, Moscow, Nauka: 53-70 (Comput. Seismol.; Iss. 8, in Russian).
- Zhidkov,M.P., and V.G.Kossobokov (1980). Identification of the sites of possible strong earthquakes - VIII: Intersections of lineaments in the east of Central Asia. In V.I.Keilis-Borok (ed.), *Earthquake Prediction and the Structure of the Earth*, Allerton Press Inc, New York: 31-44 (Comput. Seismol., volume 11).