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International Centre for Theoretical Physics**



2142-Presentation

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Development**

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Asi-Sisma Presentations

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Present-day deformation field in the Central Mediterranean region revealed by the coupled use of thermomechanical modelling and geodetic data

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GFM group

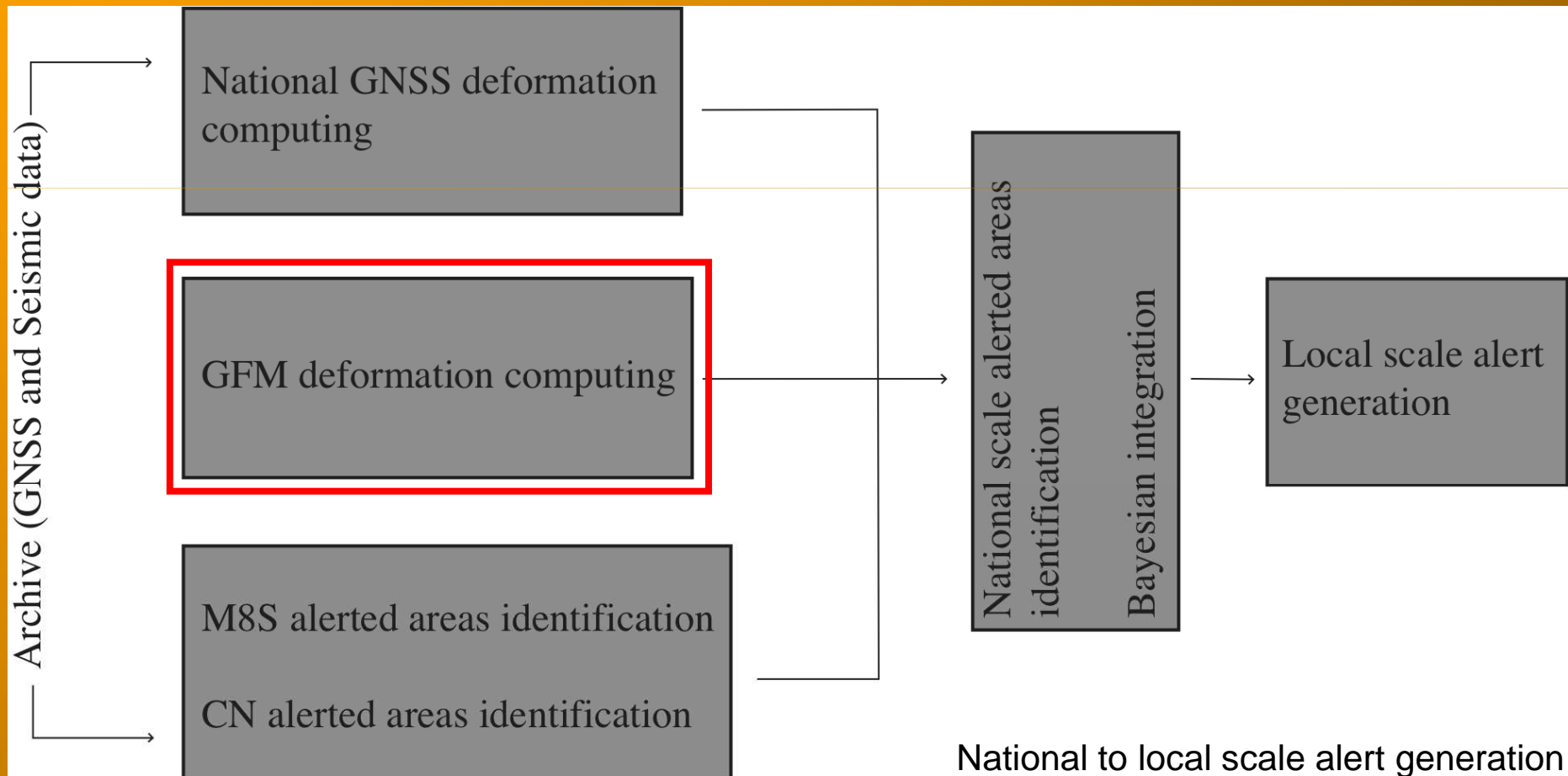


Università degli Studi di Milano

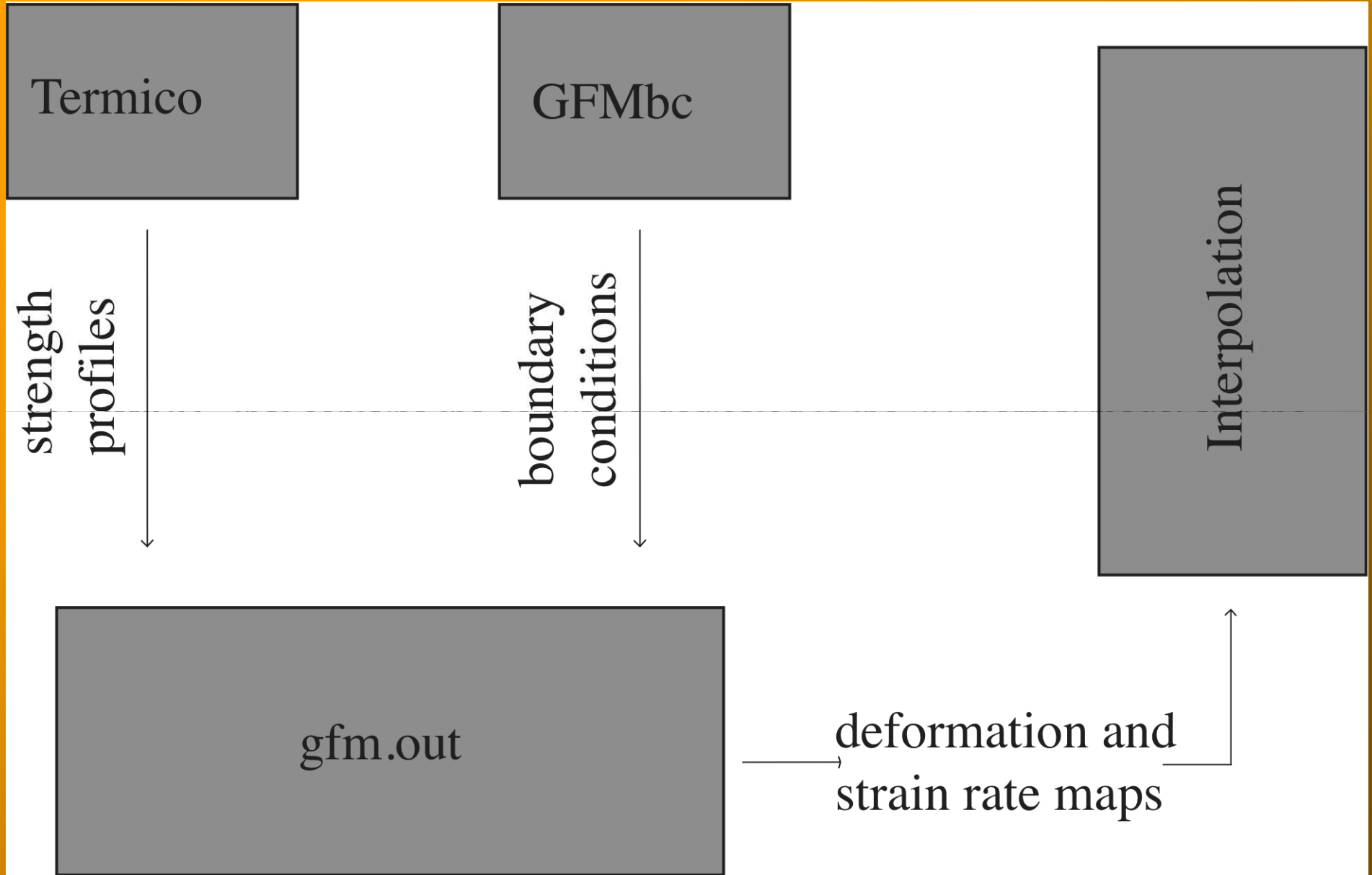
Project SISMA

SISMA (Seismic Information System for Monitoring and Alert) is a pilot project for the seismic hazard monitoring in Italy, funded by ASI (Italian Space Agency).

SISMA is a multiscale and multidisciplinary approach to the analysis of the seismic risk on the Italian Peninsula, that combines probabilistic methodologies with deterministic techniques in order to evaluate the hazard in a more consistent way.



GFM modulus



Numerical Model

Numerical model contains two different finite element model:

- *Finite element 2D tectonic model*, with “Thin Sheet” approach and spherical coordinates in which rheological heterogeneities, computed by a thermal model, are considered:
 - Tectonic velocity field (intermediate product)
 - Deformation map (GFM product)
 - Strain-rate map (GFM product)
- *Finite element 3D thermal model*:
 - Litospheric temperature vertical profiles (intermediate product)
 - Strength profiles, and effective viscosity (intermediate product)

Tectonic model

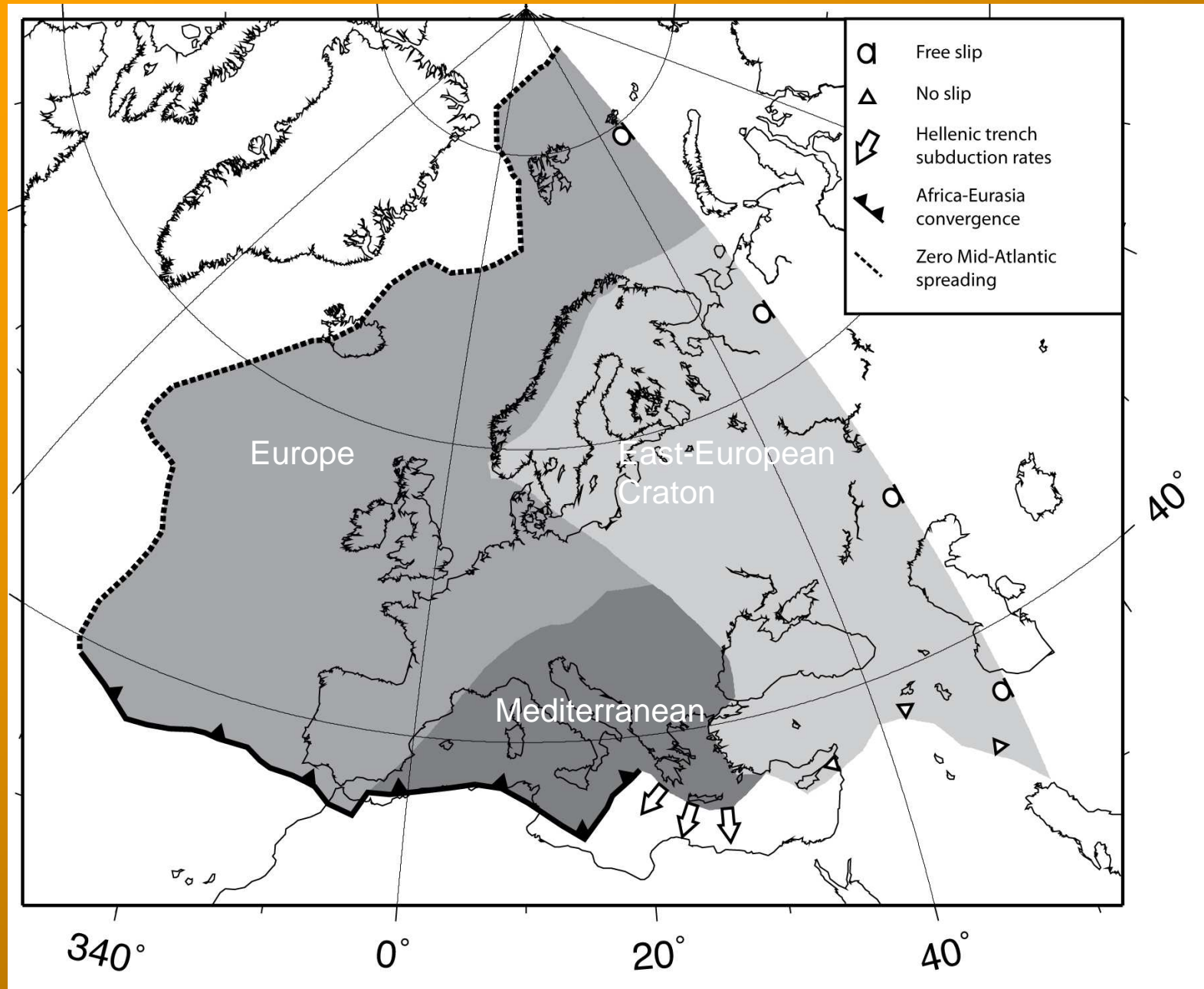
Tectonic model solves the momentum equations in spherical coordinates on a bidimensional grid, computing the horizontal velocities:

$$\begin{aligned} & \frac{\partial}{\partial \theta} \left[2\bar{\mu} \left(\frac{\partial}{\partial \theta} u_\theta - \frac{1}{2} \left(\frac{\partial u_\theta}{\partial \theta} + \frac{1}{\sin \theta} \frac{\partial u_\Phi}{\partial \Phi} + u_\theta \cot \theta \right) \right) \right] + \frac{1}{\sin \theta} \frac{\partial}{\partial \Phi} \left[\bar{\mu} \left(\frac{1}{\sin \theta} \frac{\partial}{\partial \Phi} u_\theta + \frac{\partial}{\partial \theta} u_\Phi - u_\Phi \cot \theta \right) \right] \\ & + \left[2\bar{\mu} \left(\frac{\partial}{\partial \theta} u_\theta - \frac{1}{\sin \theta} \frac{\partial}{\partial \Phi} u_\Phi - u_\theta \cot \theta \right) \right] \cot \theta = \frac{g\rho_c R}{2L} \left(1 - \frac{\rho_c}{\rho_m} \right) \frac{\partial}{\partial \theta} S^2 \end{aligned} \quad (1)$$

$$\begin{aligned} & \frac{\partial}{\partial \theta} \left[\bar{\mu} \left(\frac{1}{\sin \theta} \frac{\partial}{\partial \Phi} u_\theta + \frac{\partial}{\partial \theta} u_\Phi - u_\Phi \cot \theta \right) \right] + \\ & \frac{1}{\sin \theta} \frac{\partial}{\partial \Phi} \left[2\bar{\mu} \left(\frac{1}{\sin \theta} \frac{\partial}{\partial \Phi} u_\theta + \frac{\partial}{\partial \theta} u_\Phi - u_\Phi \cot \theta \right) \right] \cot \theta = \frac{g\rho_c R}{2L} \left(1 - \frac{\rho_c}{\rho_m} \right) \frac{1}{\sin \theta} \frac{\partial}{\partial \Phi} S^2 \end{aligned} \quad (2)$$

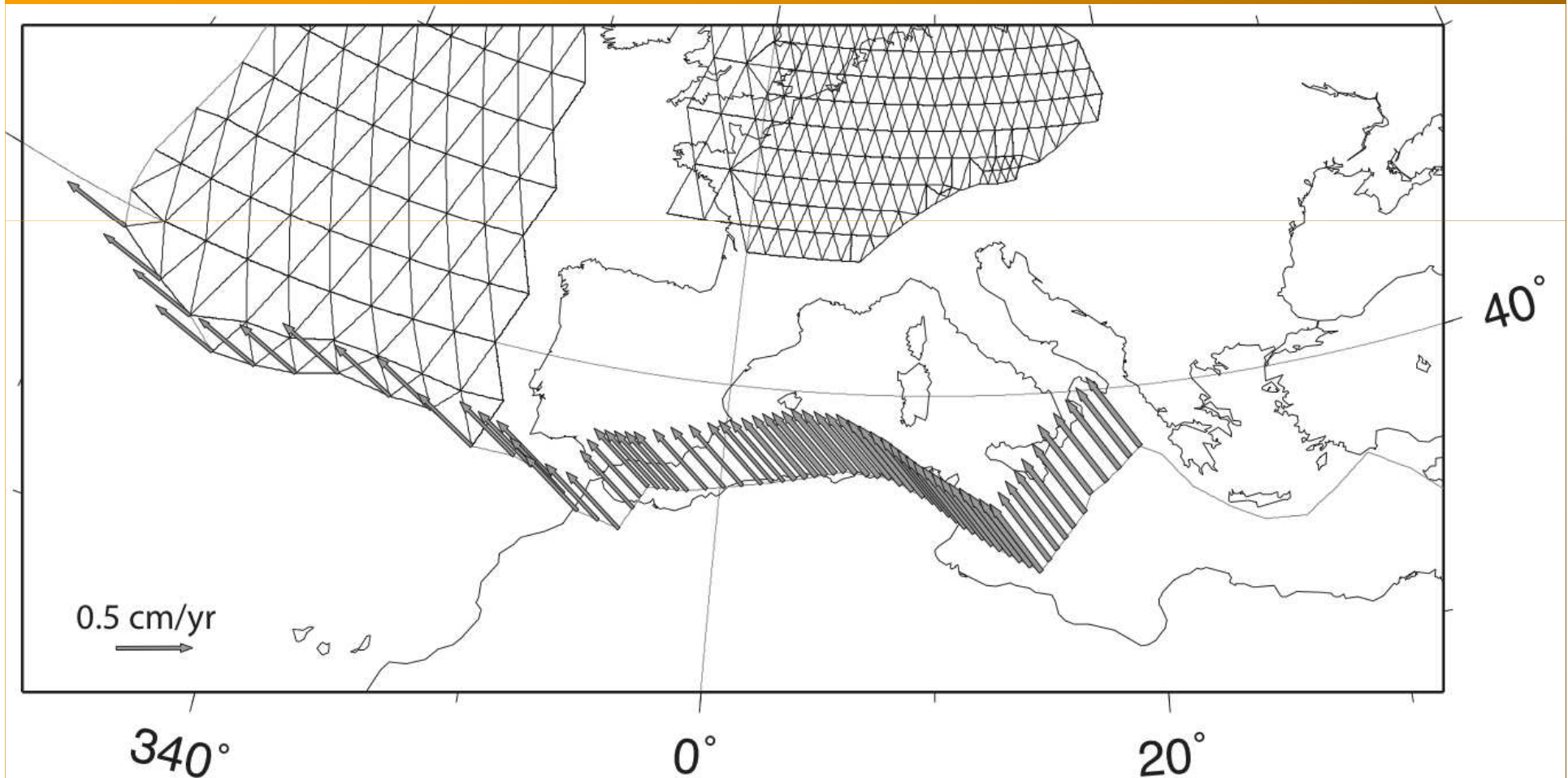
Starting from the horizontal velocity components, horizontal *strain rate* is predicted using the procedure described in Devoti et al. (2002)

Tectonic Model: Computation Domain



Tectonic model: mesh and boundary conditions

Boundary conditions (Africa-Eurasia convergence) are computed starting from ITRF 2005 solutions (Altamimi et al., 2007) through the Eulerian poles estimation (Noquet et al., 2001).



Tectonic model: boundary conditions computation

- Devoti et al. (2002)
- McClusky et al. (2003)



Tectonic model: effective viscosities

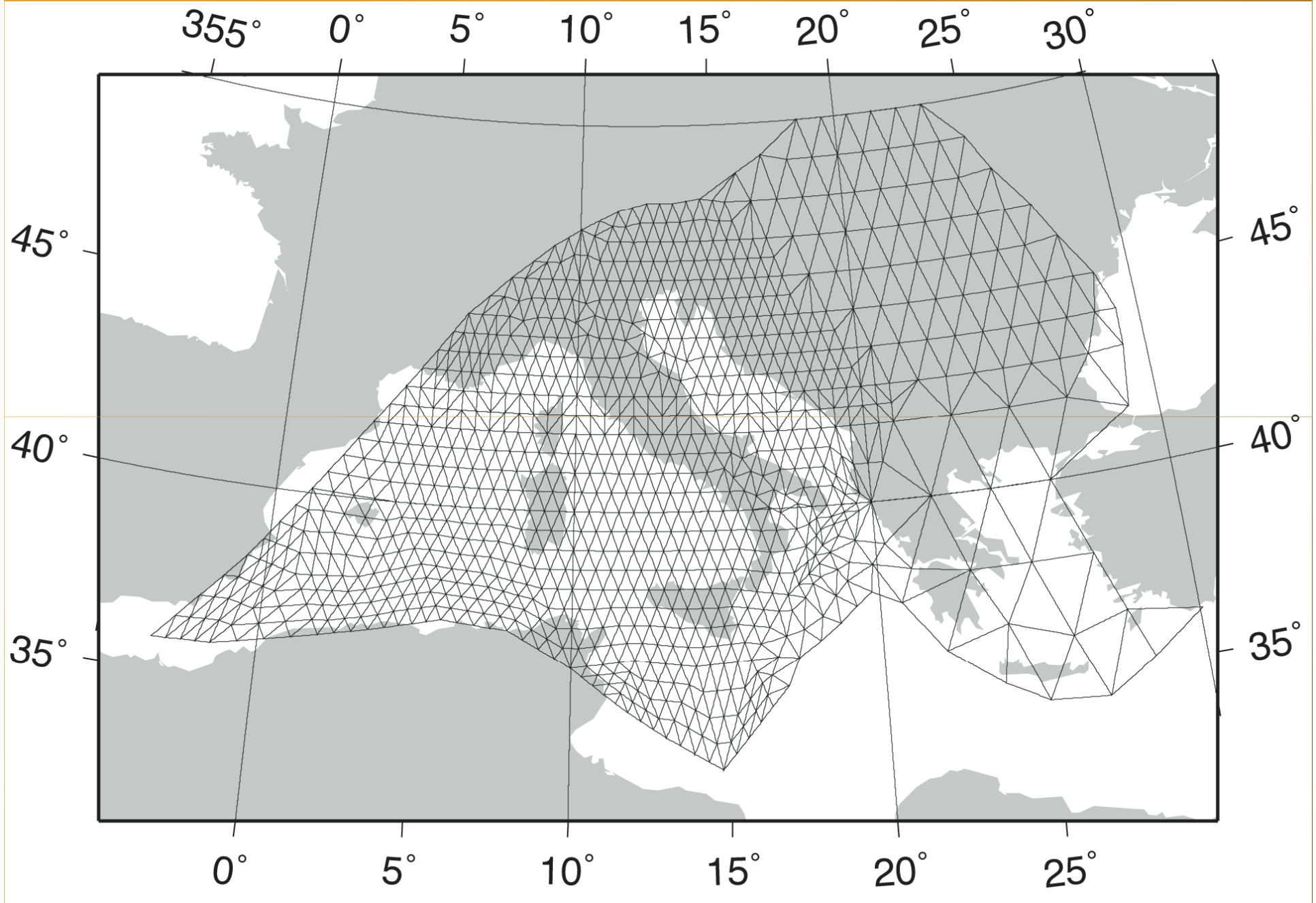
Model computation domain is made of 3 rheologically differentiated blocks, 2 with fixed effective viscosity:

- Europe: effective viscosity = 10^{25} Pas
- East-European craton: effective viscosity = 10^{27} Pas

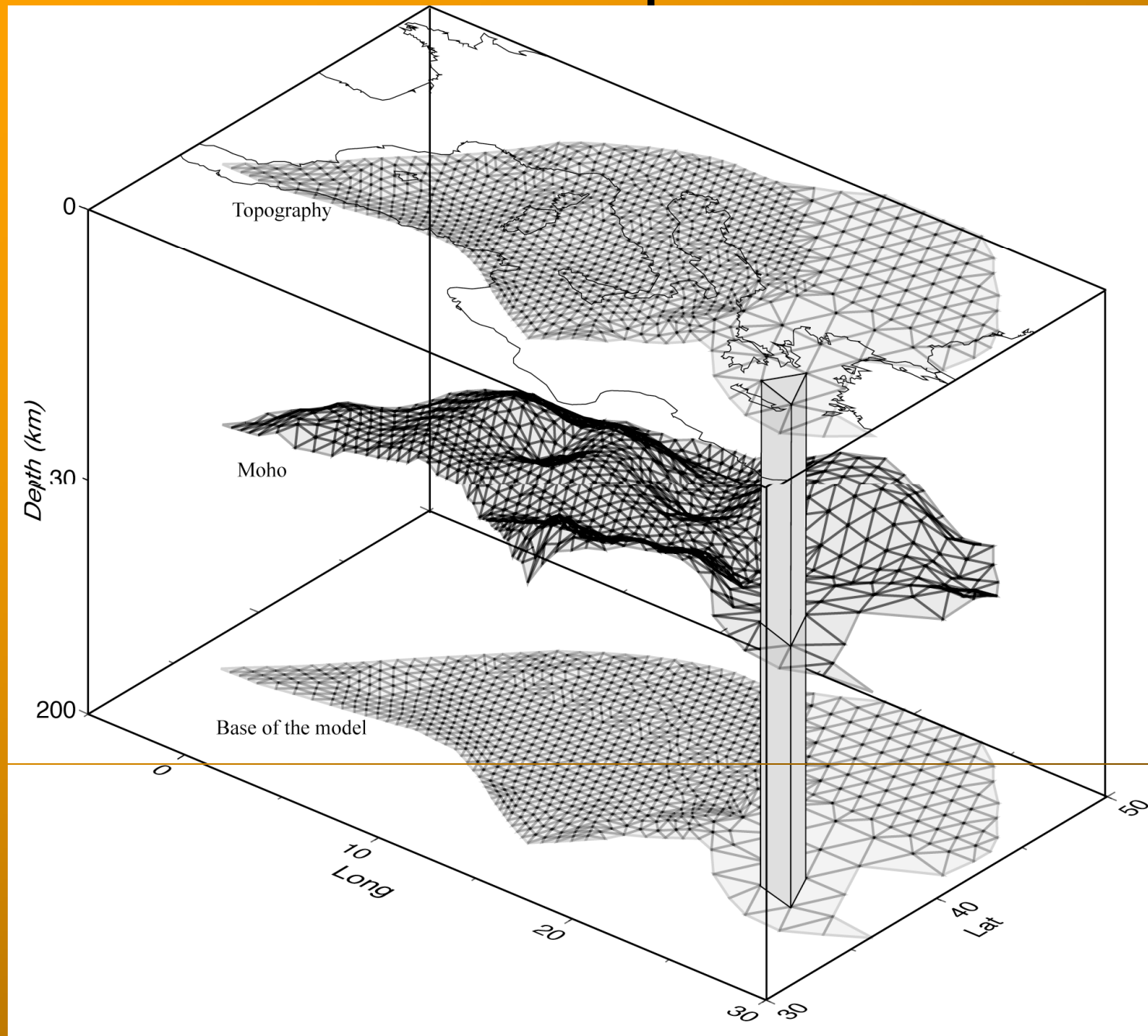
And one with calculated effective viscosity:

- Mediterranean: effective viscosity = calculated by finite element 3D thermal model coupled with a rheological analysis

Thermal model: computation domain



Thermal model: computation domain



Thermal model: equations and boundary conditions

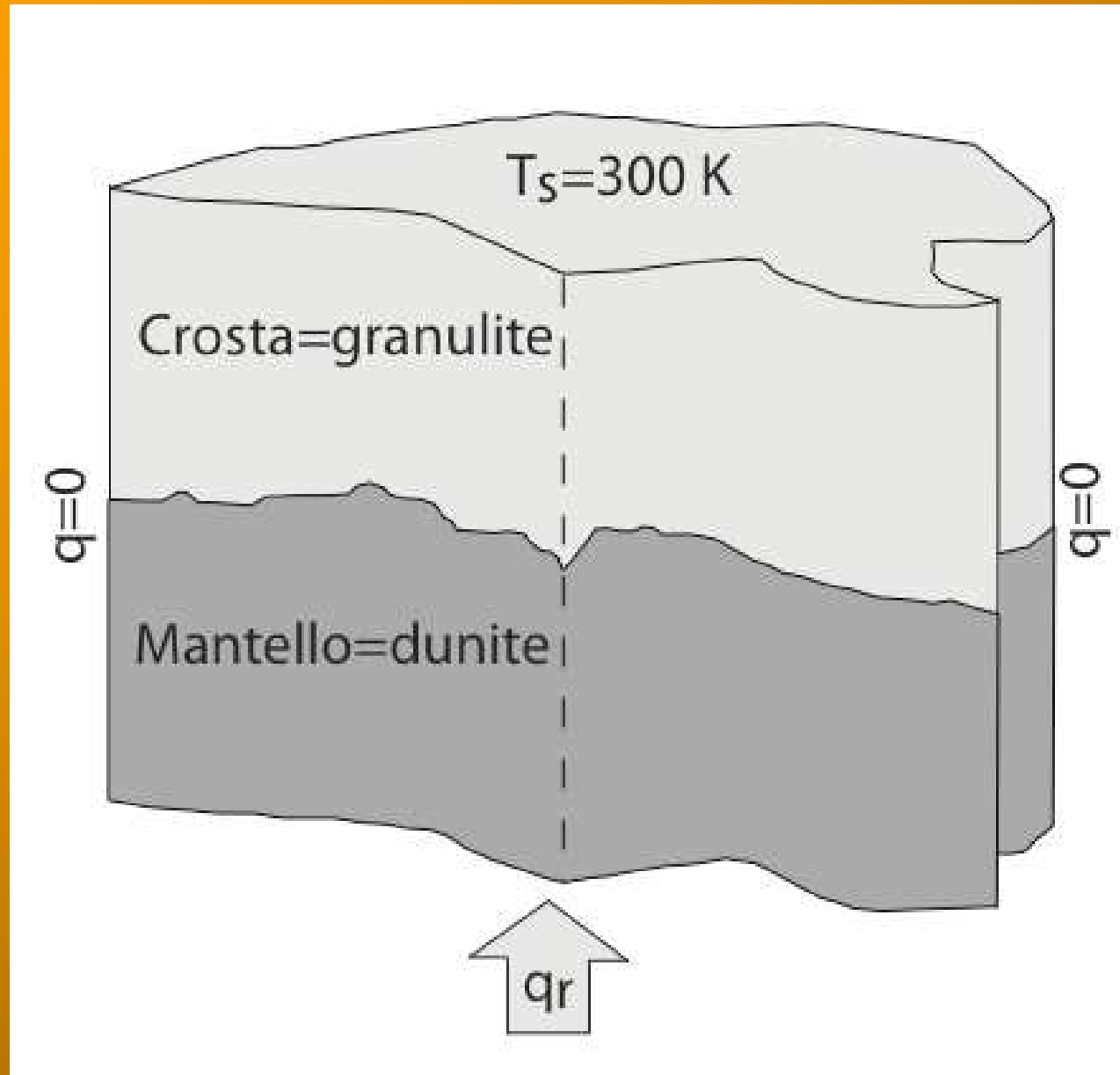
Thermal model solves the conduction equation on a 3D stratified grid, computing the vertical temperature profiles of the lithosphere:

$$\nabla \cdot (k \nabla T) + \rho H = 0$$

Model boundary conditions are:

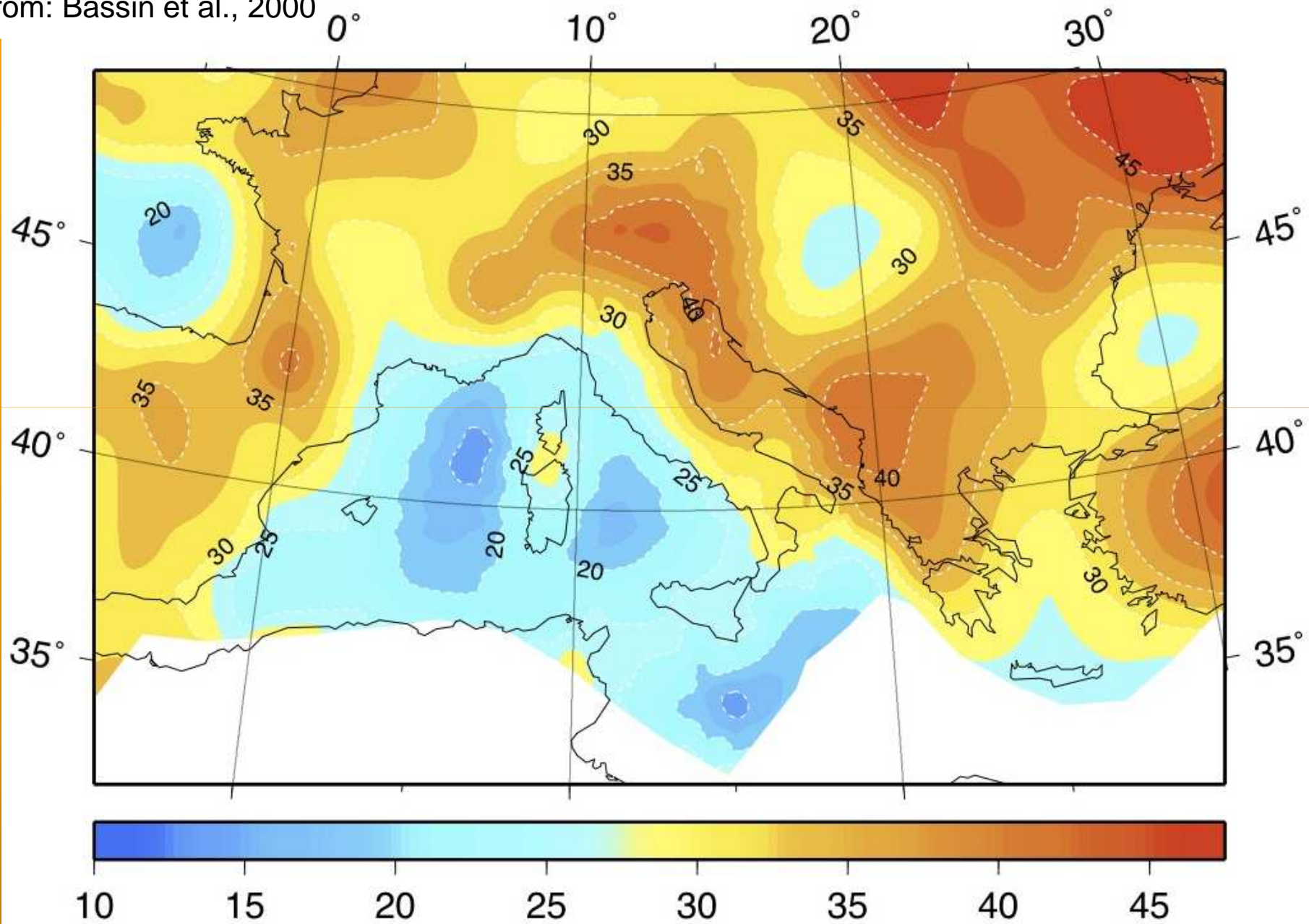
- Surface temperature of 300 K
- Zero heat flow at the lateral boundaries of the mesh
- Residual heat flow at the base of the model, calculated starting from the surface heat flow $q_r = 0.6q_s$ (Pollack & Chapman, 1977).

Thermal model: equations and boundary conditions



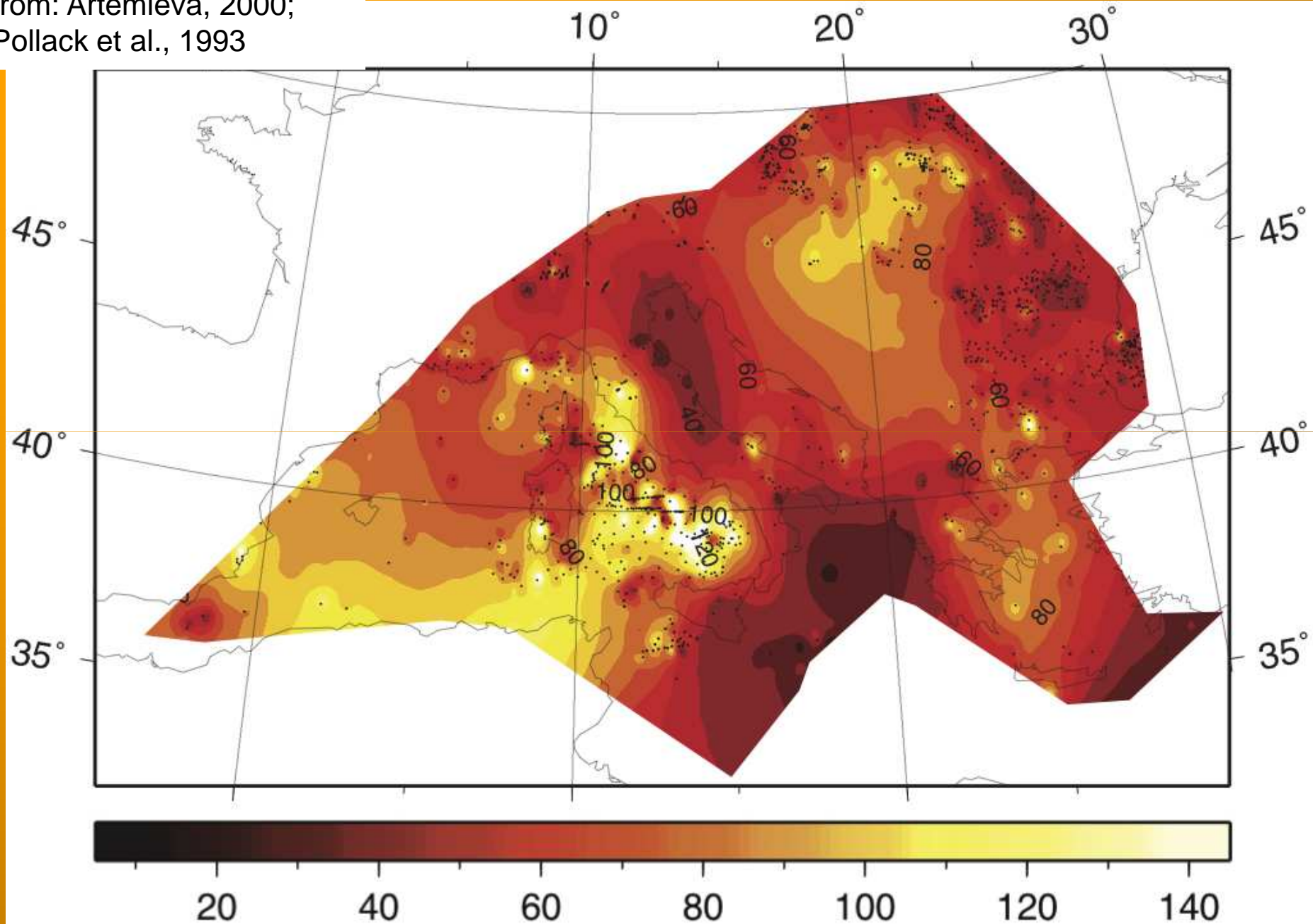
Models database: crustal thickness (km)

From: Bassin et al., 2000

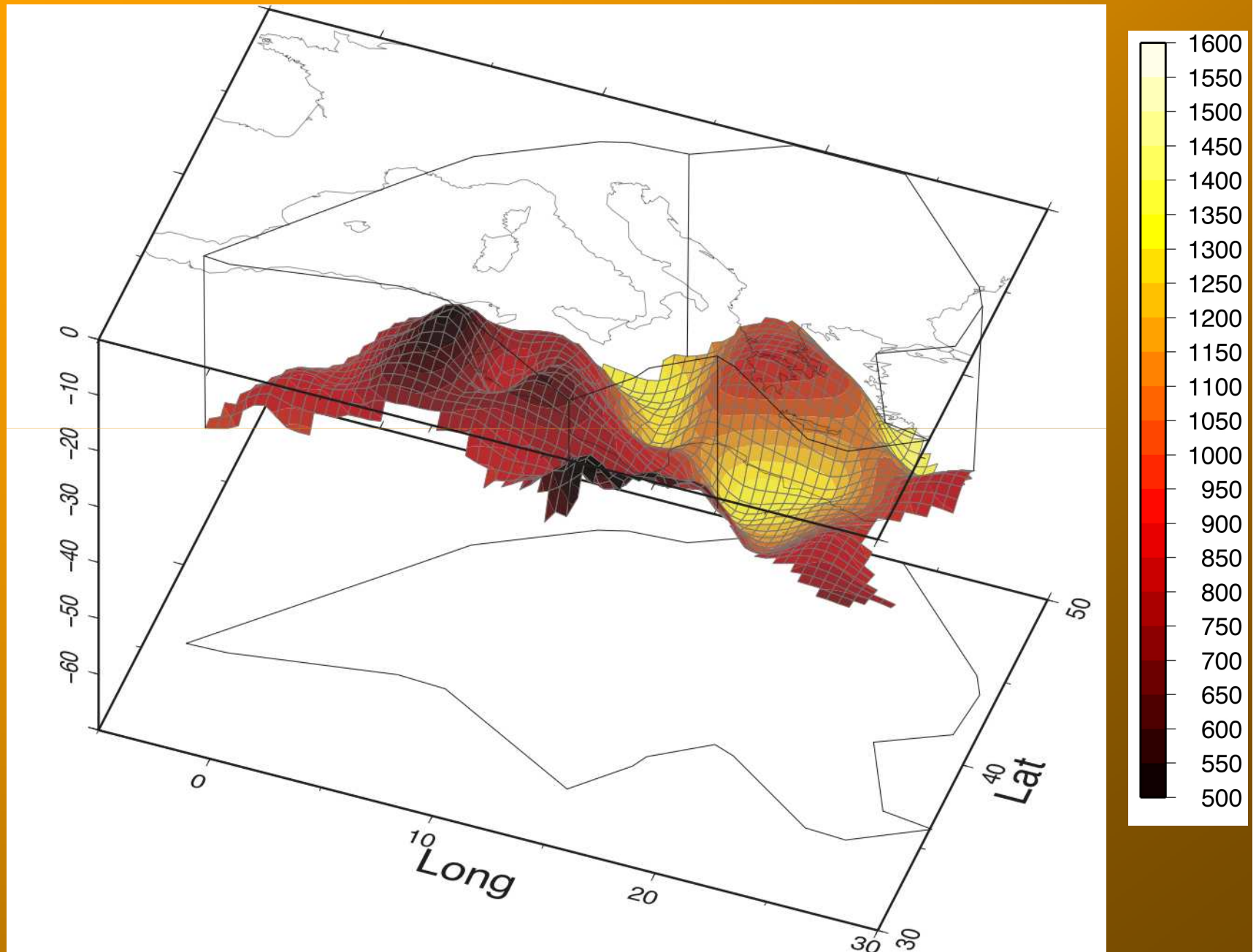


Models database: surface heat flow (mW/m²)

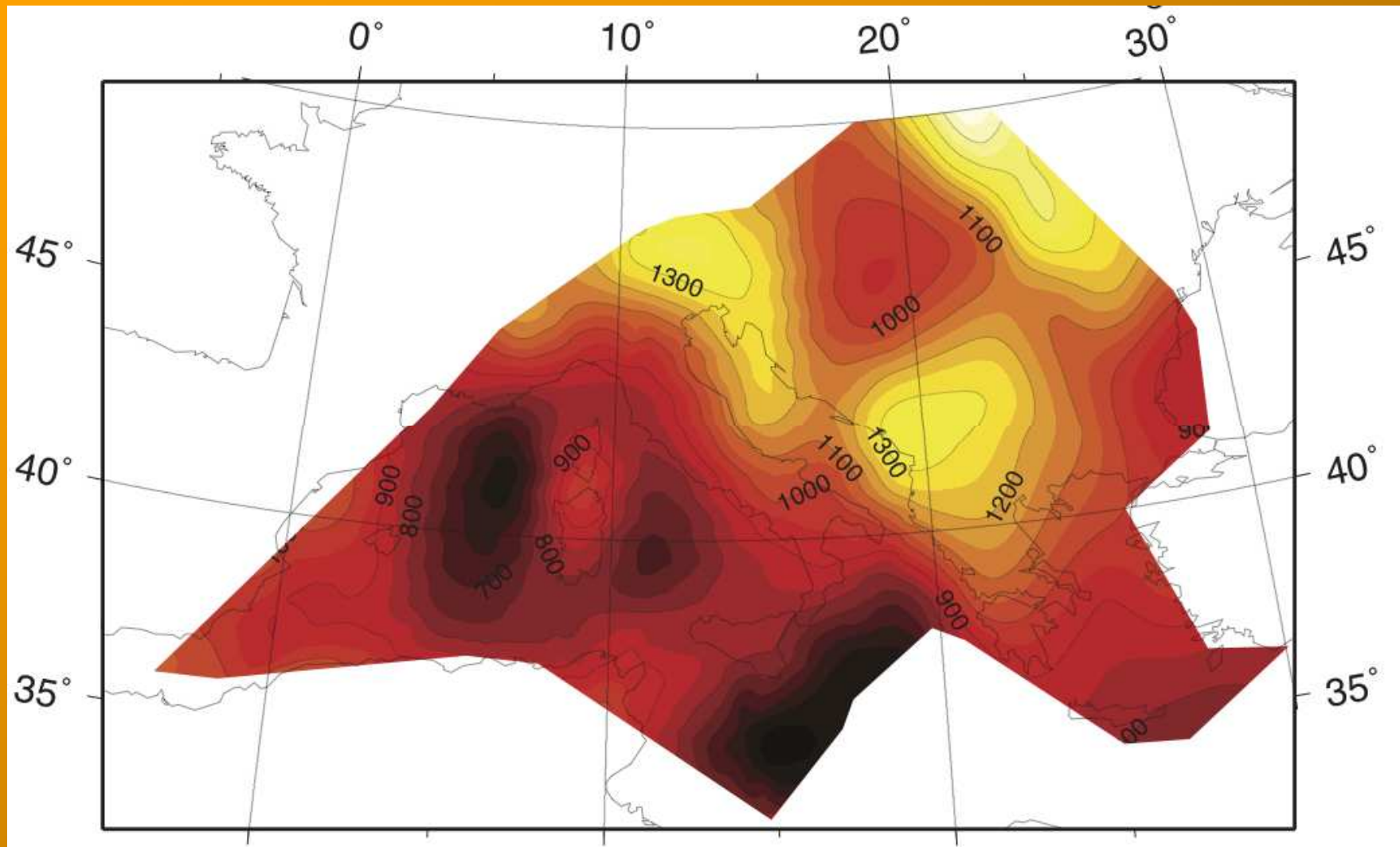
from: Artemieva, 2000;
Pollack et al., 1993



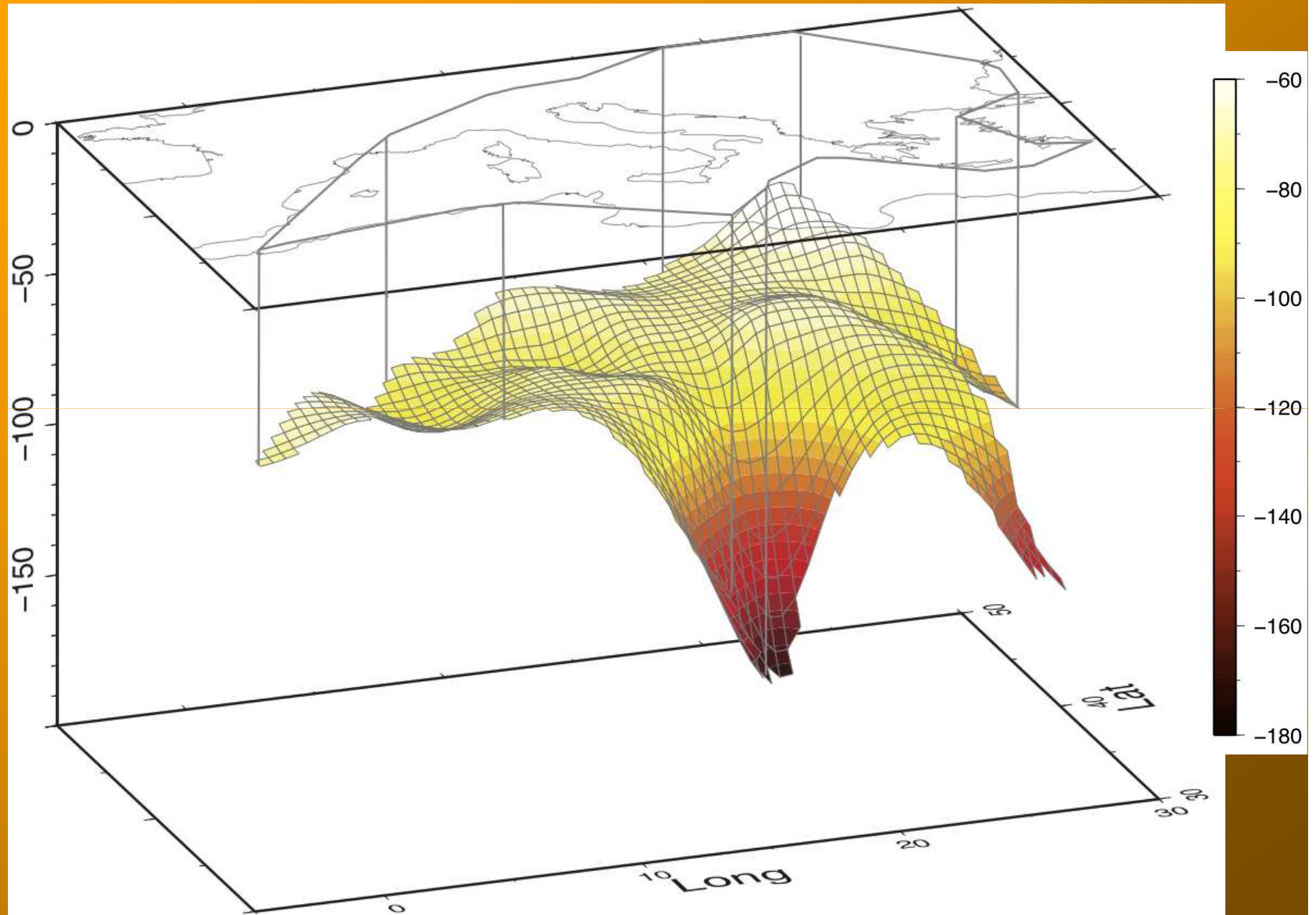
Thermal model: temperature at the Moho depth (K)



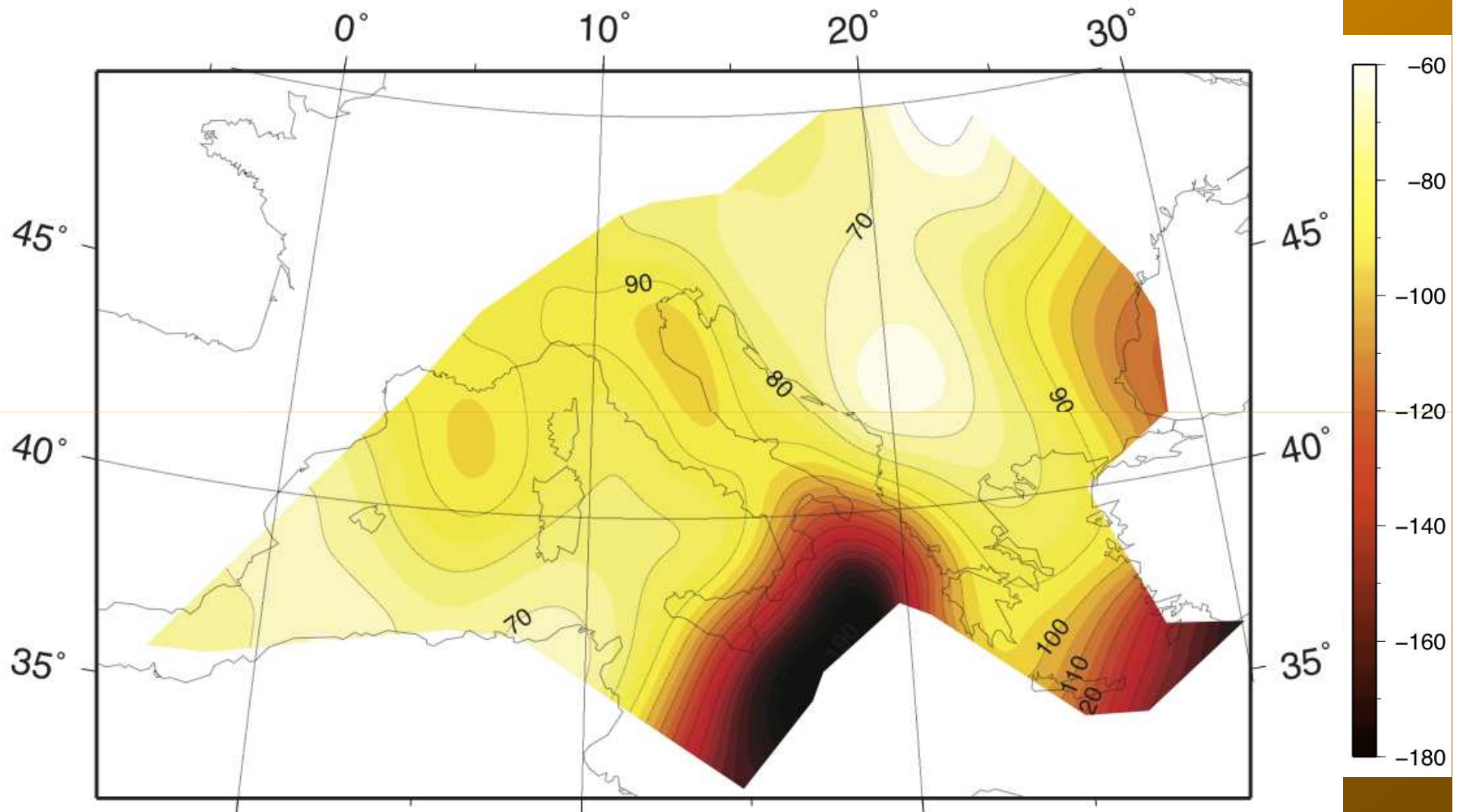
Thermal model: temperature at the Moho depth (K)



Thermal model: thermal lithosphere depth (km)



Thermal model: thermal lithosphere depth (km)



Rheological analysis

Starting from temperature profiles, strength profiles are computed:

Fragile behavior (Ranalli & Murphy, 1987):

$$\sigma_B = (\sigma_H - \sigma_V)_B = \beta \cdot r \cdot \rho \cdot g$$

With $\beta=3$ for compressive regime, $\beta=1.2$ for strike slip regime and $\beta=0.75$ for normal regime.

Ductile behavior (Weertman & Weertman, 1975):

$$\sigma_D = \left(\frac{\dot{\varepsilon}}{\dot{\varepsilon}_0} \right)^{\frac{1}{n}} \cdot \exp\left(\frac{E_a}{nRT} \right)$$

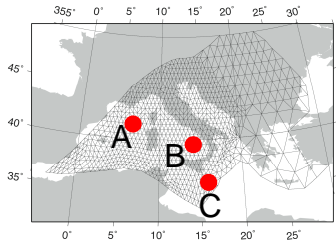
With reference strain rate $\dot{\varepsilon} = 10^{-16} \text{ s}^{-1}$

Strength profiles:

$$\sigma_y = \min\{\sigma_B, \sigma_D\}$$

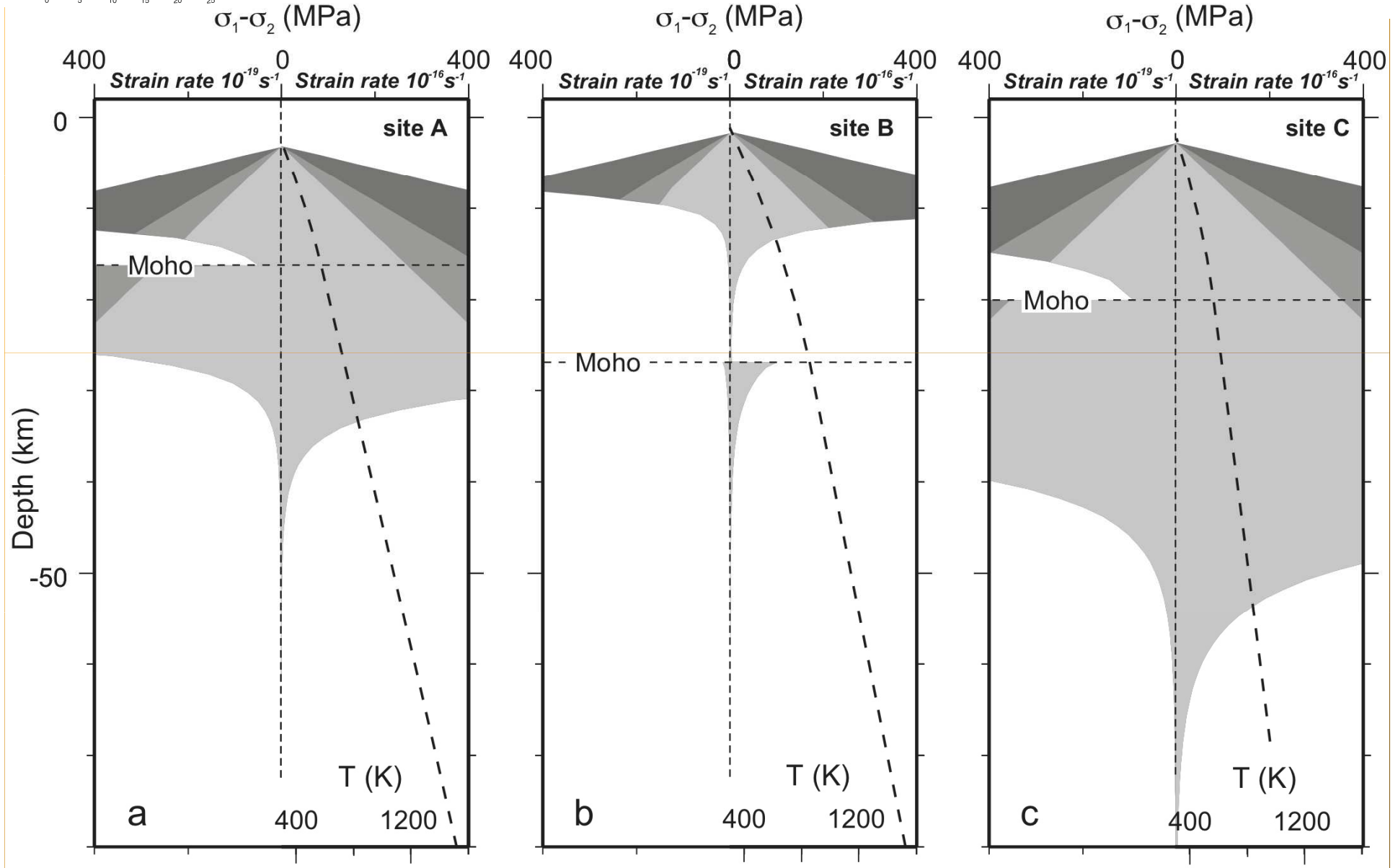
Effective viscosity:

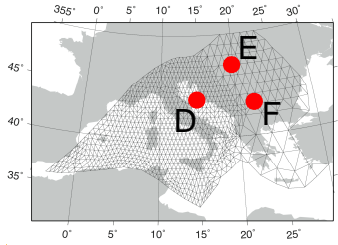
$$\mu_{eff} = \frac{1}{\dot{\varepsilon} L} \int_0^L \sigma_y dy$$



Strength profiles

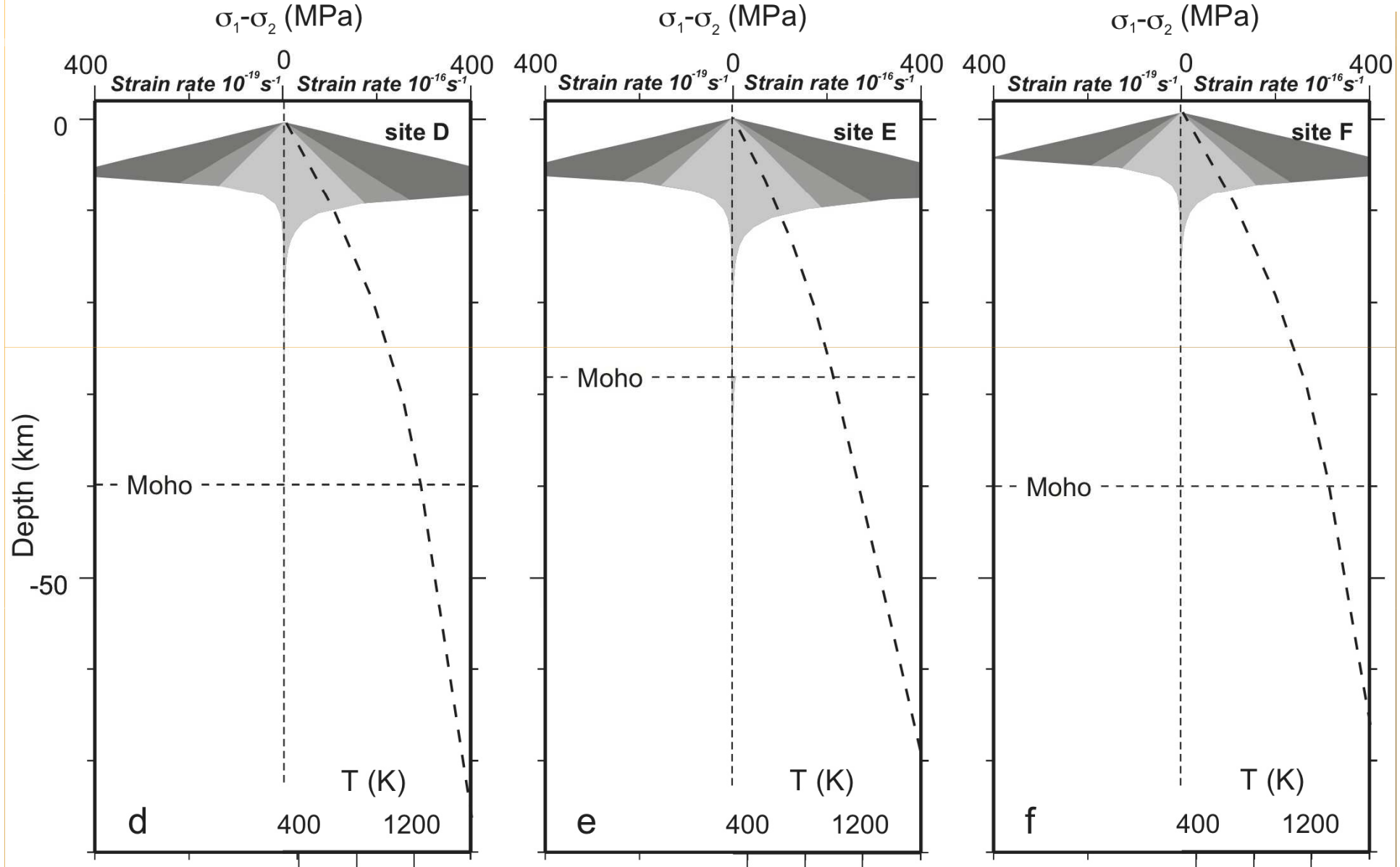
- thrust
- strike slip
- normal



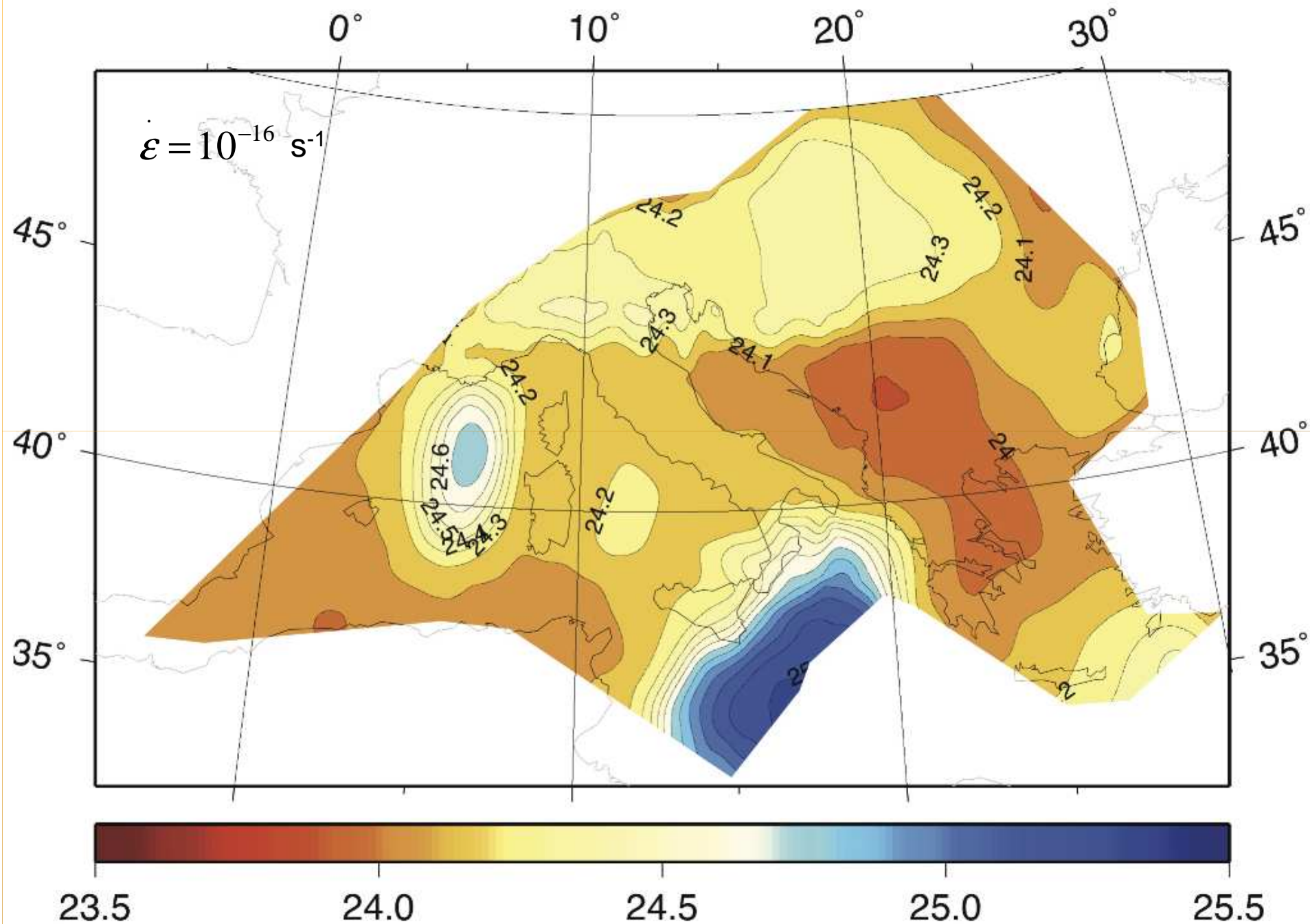


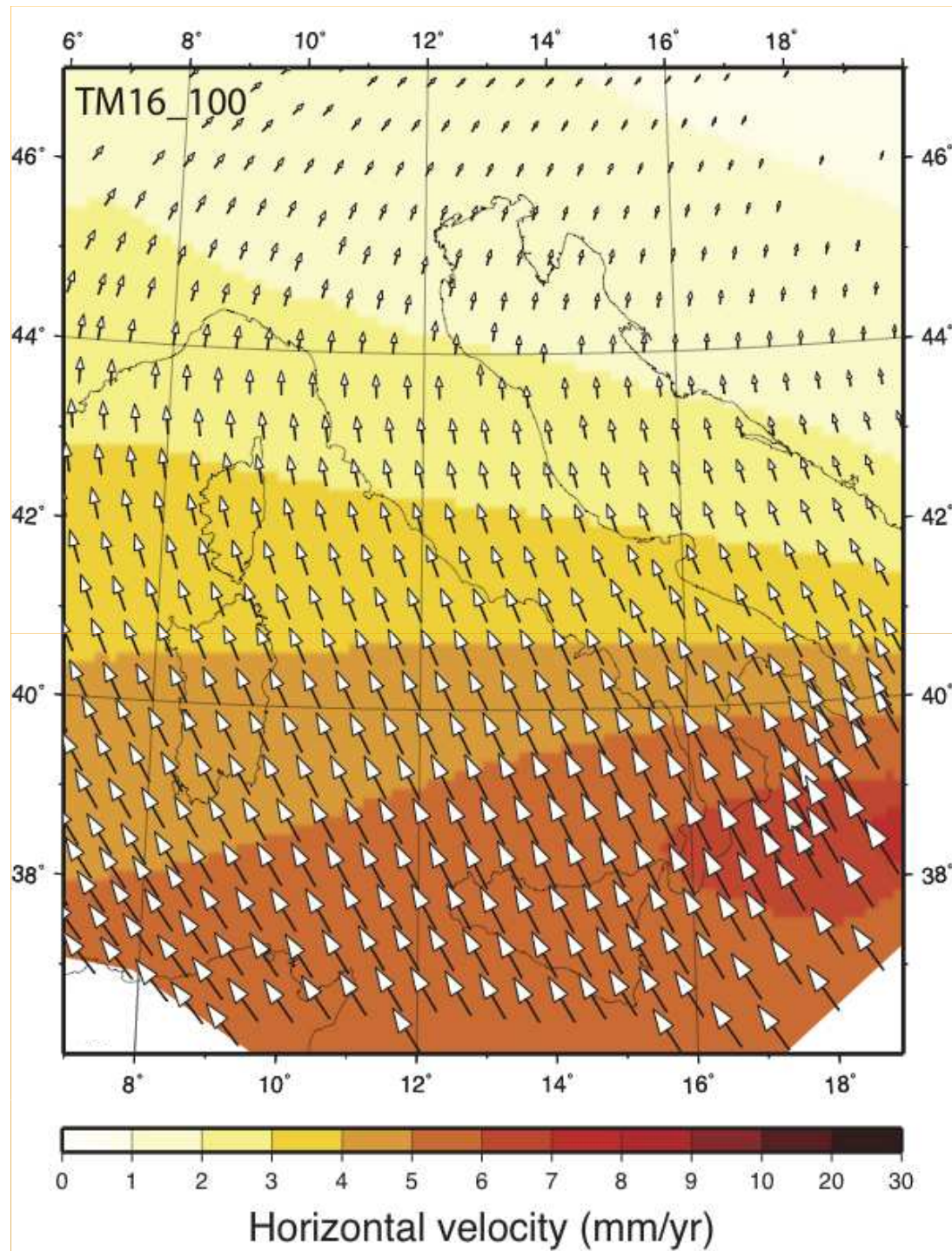
Strength profiles

- thrust
- strike slip
- normal



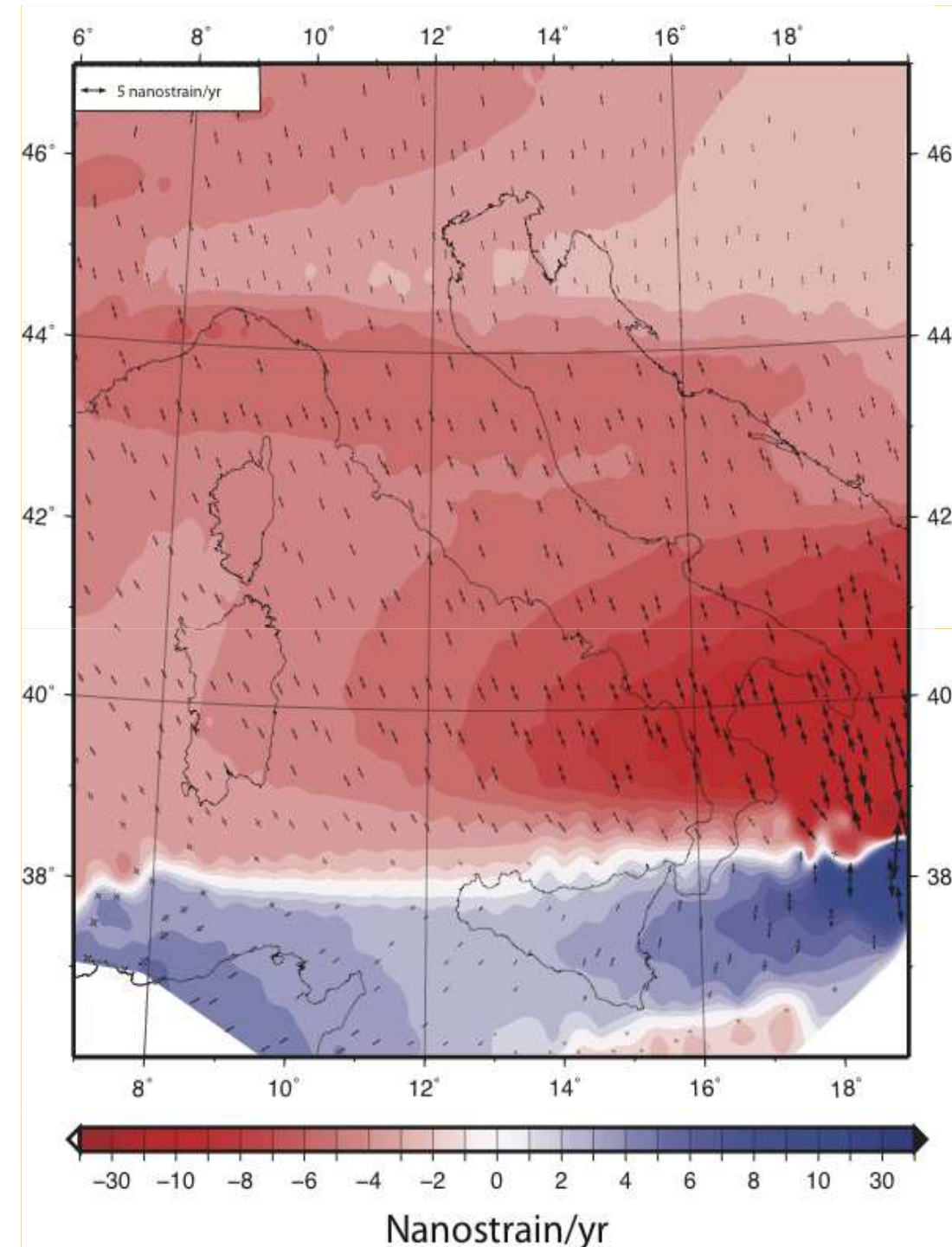
Effective viscosity ($\log_{10}\mu_{\text{eff}}$ (Pas))





Model results: velocity field

Tectonic model computes the horizontal components of velocity for each node of the reference mesh (both longitudinal and latitudinal components).



GFM products: Strain-rate field

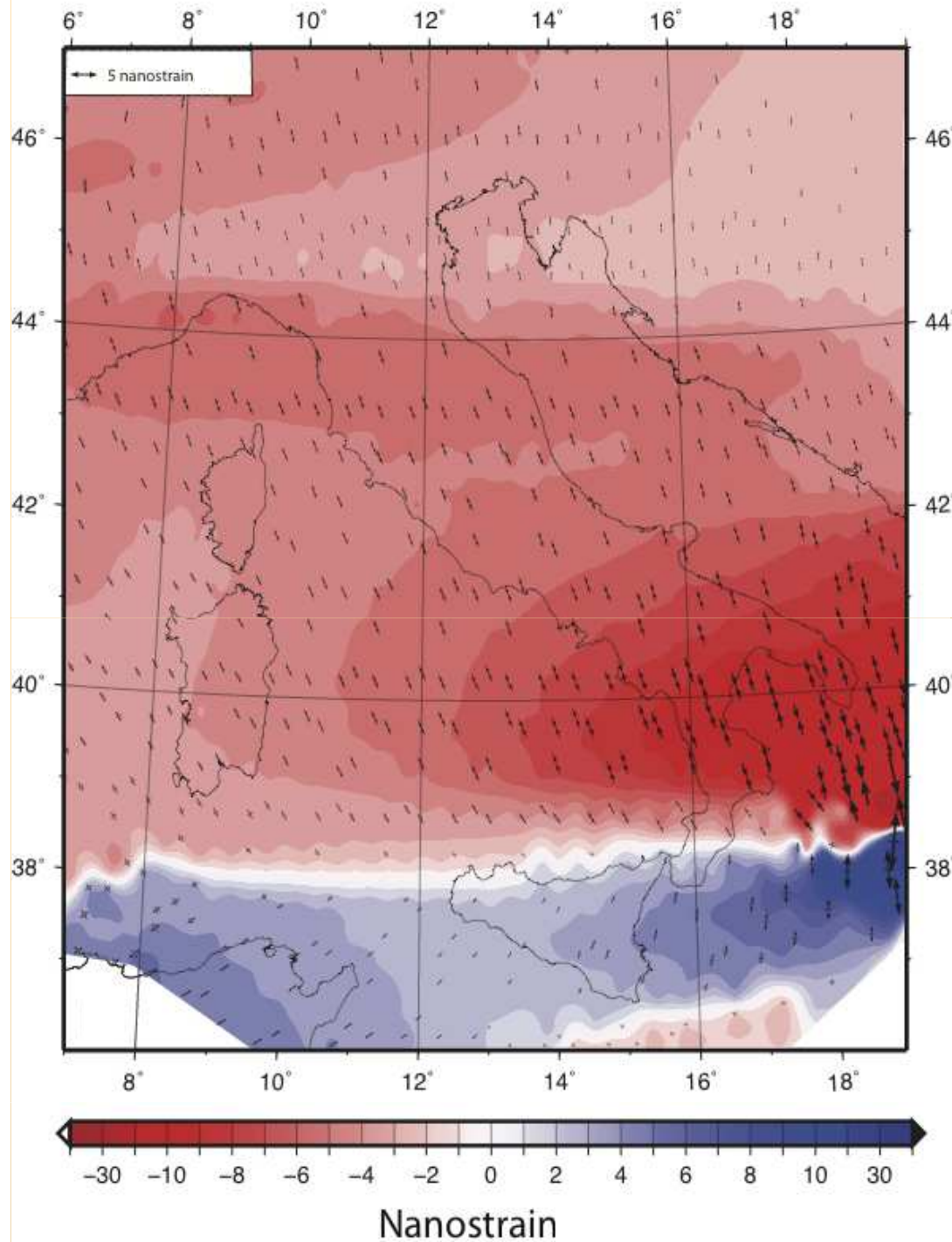
Starting from the computed velocity field, strain-rate is calculated for each element of the reference mesh, using the procedure described in Devoti et al. (2002) for triangular elements.

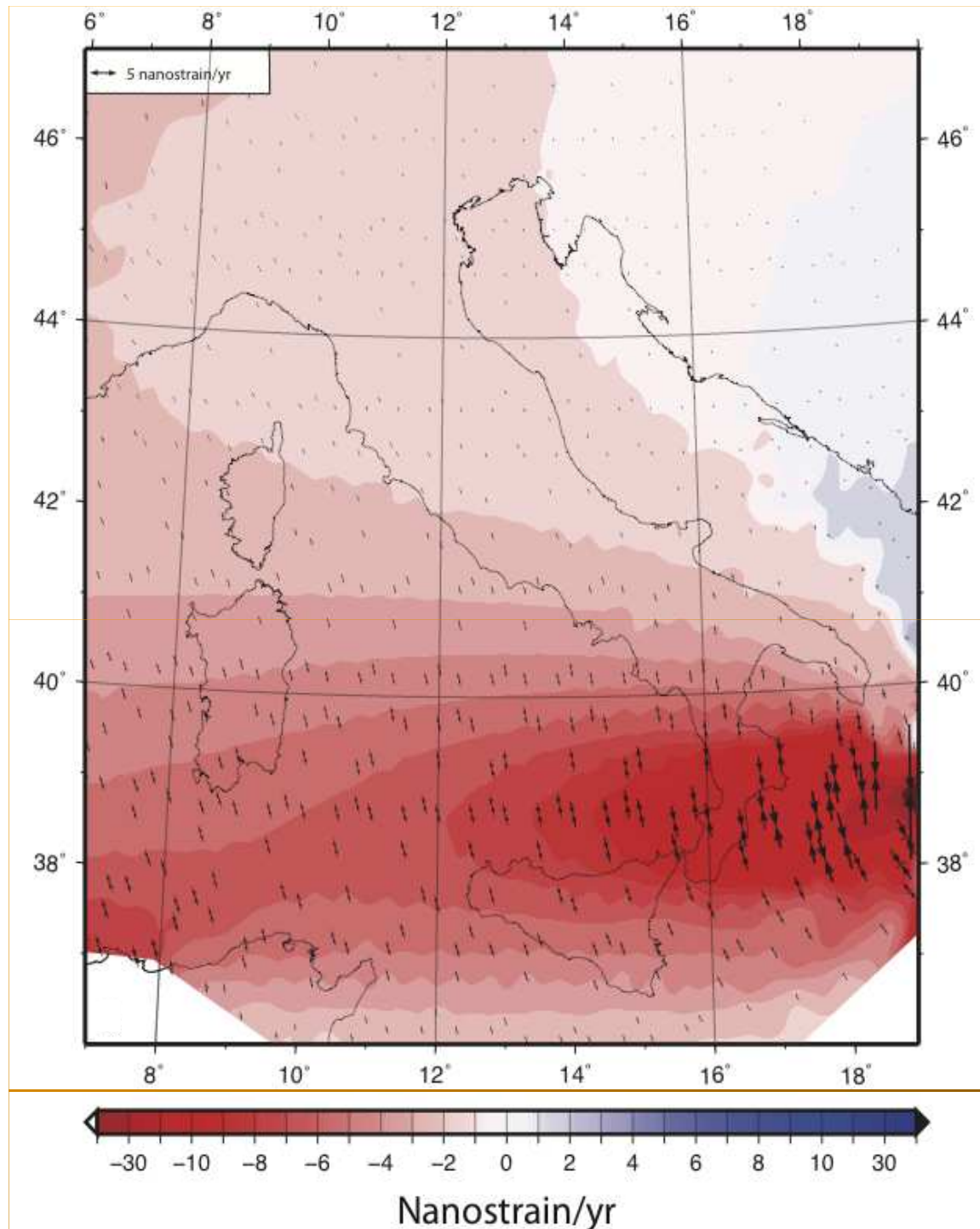
Negative strain-rate indicates compression (red), while positive strain-rate indicates extension (blue). In the map strain-rate eigenvalues are also shown (black arrows).

GFM products: Deformation field

Starting from the strain-rate field, deformation is calculated for each element of the reference mesh.

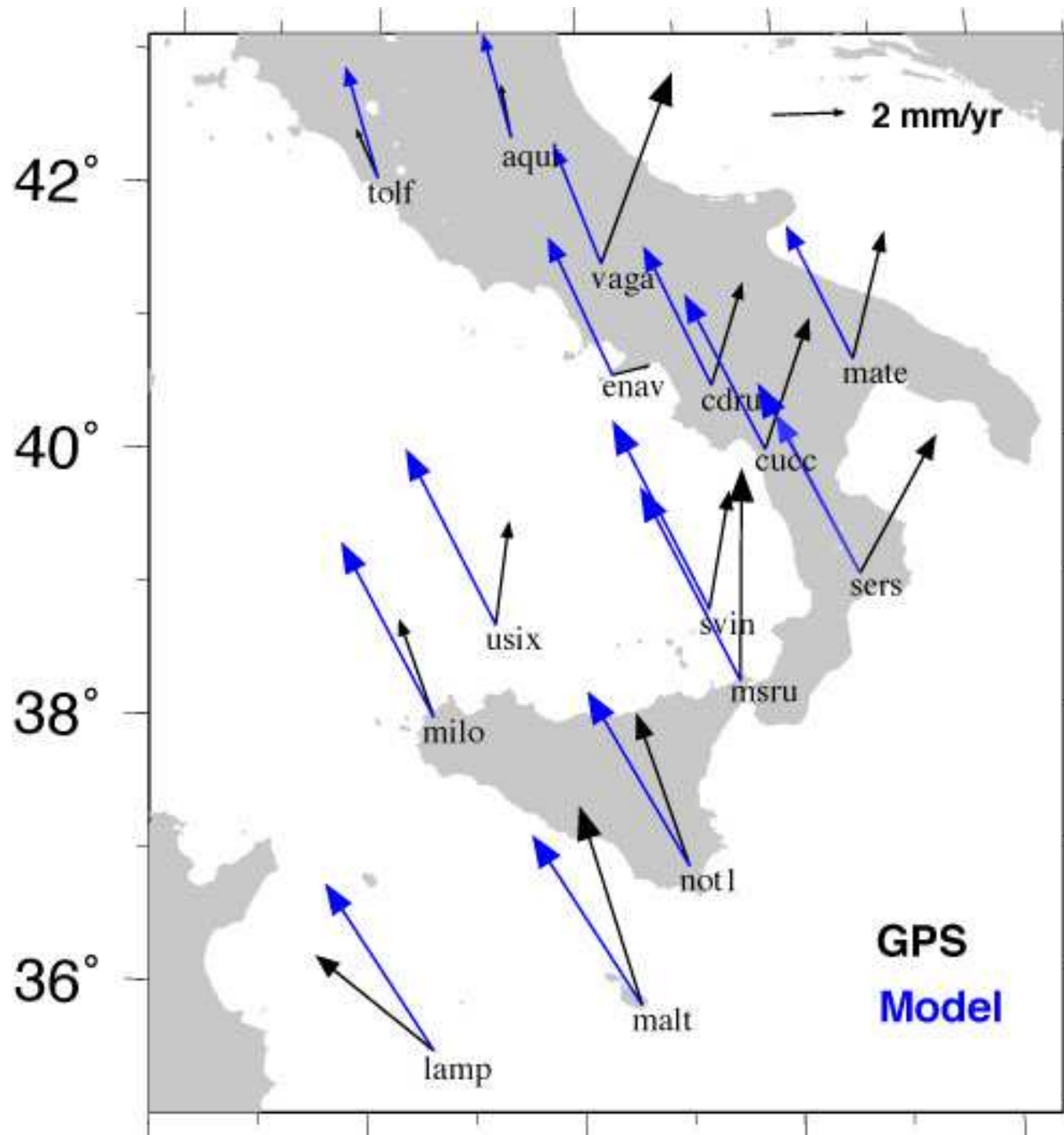
Negative deformations values indicate compression (red), while positive deformations values indicate extension (blue). In the map deformation eigenvalues are also shown (black arrows).





Tectonic deformation Fixed viscosity

Horizontal strain-rate computed by the tectonic model without considering rheological heterogeneities in the Mediterranean domain (effective viscosity is fixed = 10^{24} Pas)



Interpolation modulus

Moreover the tectonic velocities computed by the model are interpolated on the net formed by GPS permanent stations, in order to statistically compare model results with GNSS data.

Products validation strategy

Products validation procedure is based on the model stability control, namely on its capacity of tolerate boundary conditions variations (boundary velocities changes in modulus and azimuth).

Procedure:

Increasing:

- a. Modulus of the fixed velocities on the southern boundary of the model (+50%)
- b. Azimuth of the fixed velocities on the southern boundary of the model (+30°)

Model must remains stable, namely:

1. Maps must still contain real numbers
2. The order of magnitude of deformation and strain-rate values must remain 10^2 nanostrain and 10^2 nanostrain/yr, respectively

Validation procedure: BC variations

Boundary conditions (BC)

Modulus= M_0 ; Azimuth= θ_0

1958	11	0.553428	321.000000
2700	11	0.553777	321.365000
2142	11	0.554112	321.720000
2696	11	0.554457	322.075000
1976	11	0.554837	322.432000
2702	11	0.555277	322.805000
2144	11	0.555774	323.180000
2706	11	0.556280	323.552000
1977	11	0.556795	323.922000
2714	11	0.557441	324.316000
2147	11	0.558072	324.710000
2718	11	0.558732	325.098000
1978	11	0.559399	325.483000
2728	11	0.564905	325.811000
2150	11	0.562979	325.831000
2730	11	0.564683	326.005000
1979	11	0.566339	326.182000
2729	11	0.566532	325.983000
2149	11	0.566769	325.789000
2731	11	0.566992	325.597000
1961	11	0.567215	325.394000
2690	11	0.568644	325.522000
2140	11	0.570030	325.642000
2693	11	0.571380	325.776000
1937	11	0.572692	325.912000
2595	11	0.573784	326.119000
2117	11	0.574864	326.324000
2599	11	0.575901	326.533000
1938	11	0.576942	326.748000



BC Variations

Modulus= $M_0+50\%$

1958	11	0.830142	321.000000
2700	11	0.830665	321.365000
2142	11	0.831168	321.720000
2696	11	0.831685	322.073000
1976	11	0.832256	322.432000
2702	11	0.832916	322.805000
2144	11	0.833661	323.180000
2706	11	0.834420	323.552000
1977	11	0.835193	323.922000
2714	11	0.836162	324.316000
2147	11	0.837108	324.710000
2718	11	0.838098	325.098000
1978	11	0.839098	325.483000
2728	11	0.847357	325.811000
2150	11	0.844469	325.831000
2730	11	0.847025	326.005000
1979	11	0.849508	326.182000
2729	11	0.849798	325.983000
2149	11	0.850154	325.789000
2731	11	0.850488	325.597000
1961	11	0.850823	325.394000
2690	11	0.852966	325.522000
2140	11	0.855045	325.642000
2693	11	0.857070	325.776000
1937	11	0.859038	325.912000
2595	11	0.860676	326.119000
2117	11	0.862296	326.324000
2599	11	0.863851	326.533000
1938	11	0.865413	326.748000



BC Variations

Modulus= $M_0+50\%$ +Azimuth= θ_0+30°

1958	11	0.830142	351.000000
2700	11	0.830665	351.365000
2142	11	0.831168	351.720000
2696	11	0.831685	352.073000
1976	11	0.832256	352.432000
2702	11	0.832916	352.805000
2144	11	0.833661	353.180000
2706	11	0.834420	353.552000
1977	11	0.835193	353.922000
2714	11	0.836162	354.316000
2147	11	0.837108	354.710000
2718	11	0.838098	355.098000
1978	11	0.839098	355.483000
2728	11	0.847357	355.811000
2150	11	0.844469	355.831000
2730	11	0.847025	356.005000
1979	11	0.849508	356.182000
2729	11	0.849798	355.983000
2149	11	0.850154	355.789000
2731	11	0.850488	355.597000
1961	11	0.850823	355.394000
2690	11	0.852966	355.522000
2140	11	0.855045	355.642000
2693	11	0.857070	355.776000
1937	11	0.859038	355.912000
2595	11	0.860676	356.119000
2117	11	0.862296	356.324000
2599	11	0.863851	356.533000
1938	11	0.865413	356.748000

Validation procedure: Surface strain-rate map

Boundary Conditions (BC)

Modulus= M_0 ; Azimuth= θ_0

LONG	LAT	S1	S2	AZIZ
9.26	74.7	0.249	-0.277E-01	-27.4
9.90	74.3	0.354	-0.773E-01	-25.0
10.5	74.7	0.225	-0.606E-01	-22.7
11.0	74.3	0.255	-0.810E-01	-23.8
11.5	74.7	0.183	-0.716E-01	-21.6
12.0	74.3	0.240	-0.989E-01	-22.7
12.5	74.7	0.171	-0.894E-01	-19.8
13.0	74.3	0.201	-0.110	-21.5
13.5	74.7	0.161	-0.104	-19.2
14.0	74.3	0.165	-0.116	-20.7
14.5	74.7	0.154	-0.114	-19.3
15.0	74.3	0.134	-0.117	-20.6
15.5	74.7	0.148	-0.119	-20.6
16.0	74.3	0.153	-0.131	-22.0
16.5	74.7	0.113	-0.124	-19.1
17.0	74.3	0.954E-01	-0.126	-20.3
17.5	74.7	0.108	-0.129	-20.3
18.0	74.3	0.114	-0.140	-21.8
18.5	74.7	0.748E-01	-0.133	-18.2
19.0	74.3	0.817E-01	-0.145	-20.2
19.5	74.7	0.945E-01	-0.147	-20.2
20.0	74.3	0.552E-01	-0.138	-20.3
20.5	74.7	0.665E-01	-0.140	-20.3
21.0	74.3	0.737E-01	-0.152	-22.0
21.5	74.7	0.596E-01	-0.149	-20.2
22.0	74.3	0.453E-01	-0.150	-21.0
22.5	74.7	0.309E-01	-0.147	-18.5
23.0	74.3	0.393E-01	-0.159	-20.6
23.5	74.7	0.497E-01	-0.161	-20.9

BC Variations

Modulus= $M_0+50\%$

LONG	LAT	S1	S2	AZIZ
9.26	74.7	0.369	-0.407E-01	-27.4
9.90	74.3	0.520	-0.114	-24.9
10.5	74.7	0.398	-0.906E-01	-22.7
11.0	74.3	0.407	-0.129	-23.6
11.5	74.7	0.288	-0.113	-21.2
12.0	74.3	0.398	-0.142	-22.5
12.5	74.7	0.279	-0.194	-20.4
13.0	74.3	0.316	-0.164	-22.0
13.5	74.7	0.257	-0.156	-19.9
14.0	74.3	0.251	-0.169	-21.3
14.5	74.7	0.219	-0.164	-19.5
15.0	74.3	0.214	-0.177	-21.0
15.5	74.7	0.208	-0.177	-20.1
16.0	74.3	0.204	-0.189	-21.5
16.5	74.7	0.171	-0.184	-19.4
17.0	74.3	0.168	-0.197	-21.0
17.5	74.7	0.161	-0.196	-20.0
18.0	74.3	0.158	-0.208	-21.3
18.5	74.7	0.126	-0.202	-18.9
19.0	74.3	0.124	-0.214	-20.5
19.5	74.7	0.116	-0.213	-19.3
20.0	74.3	0.927E-01	-0.215	-20.2
20.5	74.7	0.111	-0.218	-20.2
21.0	74.3	0.108	-0.231	-21.7
21.5	74.7	0.753E-01	-0.224	-19.0
22.0	74.3	0.546E-01	-0.225	-19.8
22.5	74.7	0.706E-01	-0.228	-19.8
23.0	74.3	0.708E-01	-0.240	-21.4
23.5	74.7	0.617E-01	-0.239	-20.1

BC Variations

Modulus= $M_0+50\%$ +Azimuth= θ_0+30°

LONG	LAT	S1	S2	AZIZ
9.26	74.7	0.346	-0.320E-01	-25.9
9.90	74.3	0.493	-0.900E-01	-23.2
10.5	74.7	0.286	-0.695E-01	-20.9
11.0	74.3	0.365	-0.102	-21.6
11.5	74.7	0.299	-0.960E-01	-20.2
12.0	74.3	0.323	-0.115	-21.2
12.5	74.7	0.257	-0.108	-19.4
13.0	74.3	0.307	-0.135	-20.7
13.5	74.7	0.212	-0.125	-17.8
14.0	74.3	0.237	-0.144	-19.3
14.5	74.7	0.231	-0.144	-18.3
15.0	74.3	0.203	-0.148	-19.6
15.5	74.7	0.193	-0.148	-18.7
16.0	74.3	0.193	-0.159	-20.1
16.5	74.7	0.156	-0.154	-17.8
17.0	74.3	0.157	-0.166	-19.4
17.5	74.7	0.146	-0.166	-18.3
18.0	74.3	0.147	-0.177	-19.8
18.5	74.7	0.139	-0.176	-18.5
19.0	74.3	0.115	-0.179	-19.6
19.5	74.7	0.104	-0.178	-18.3
20.0	74.3	0.106	-0.189	-19.9
20.5	74.7	0.962E-01	-0.187	-18.4
21.0	74.3	0.749E-01	-0.190	-19.4
21.5	74.7	0.893E-01	-0.193	-19.6
22.0	74.3	0.917E-01	-0.204	-21.1
22.5	74.7	0.558E-01	-0.198	-17.8
23.0	74.3	0.367E-01	-0.200	-18.7
23.5	74.7	0.753E-01	-0.208	-20.7

Thank you for your attention...