



**The Abdus Salam  
International Centre for Theoretical Physics**



**2240-27**

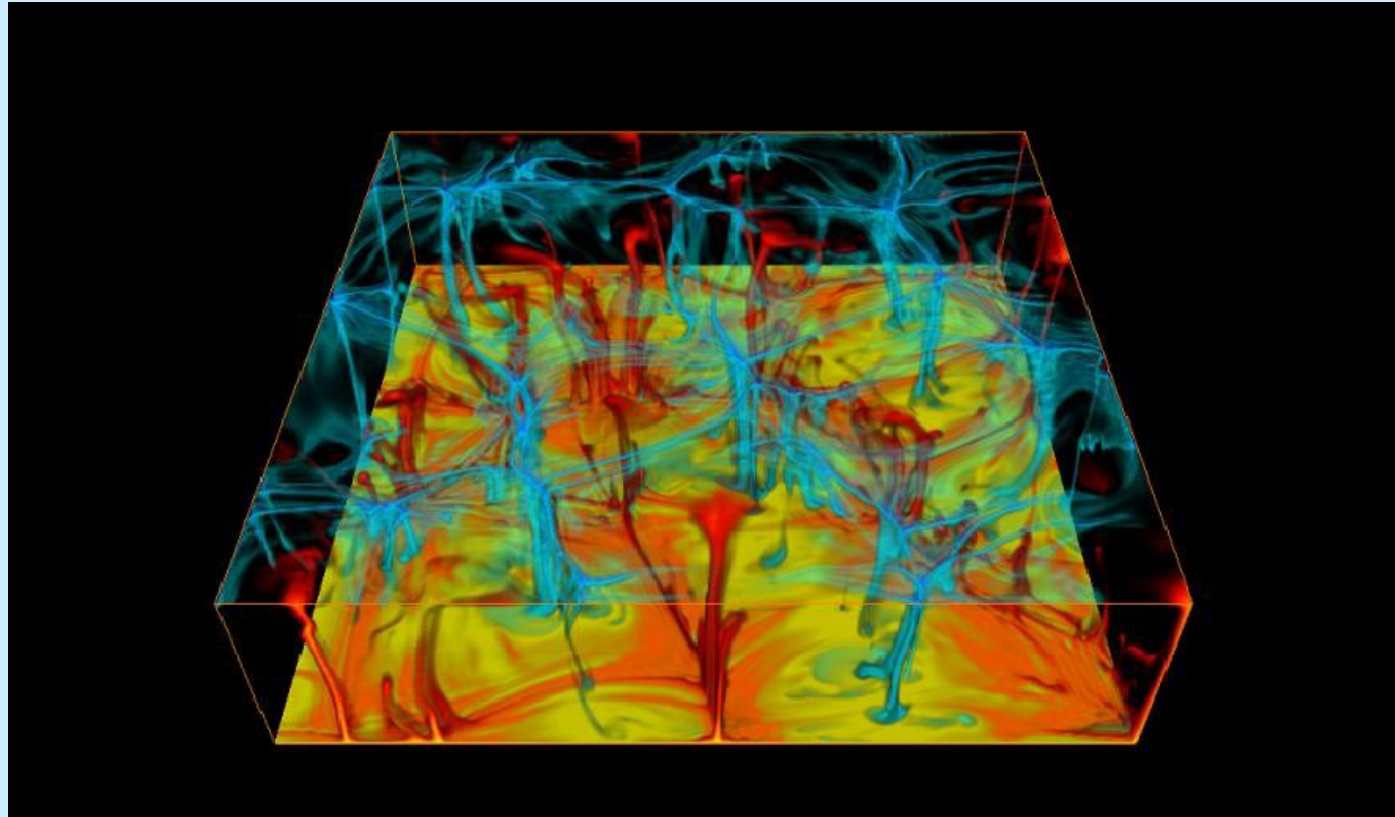
**Advanced School on Scaling Laws in Geophysics: Mechanical and  
Thermal Processes in Geodynamics**

*23 May - 3 June, 2011*

**Convection - Part II**

Claude JAUPART

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France*



**$Ra = 10^8$**

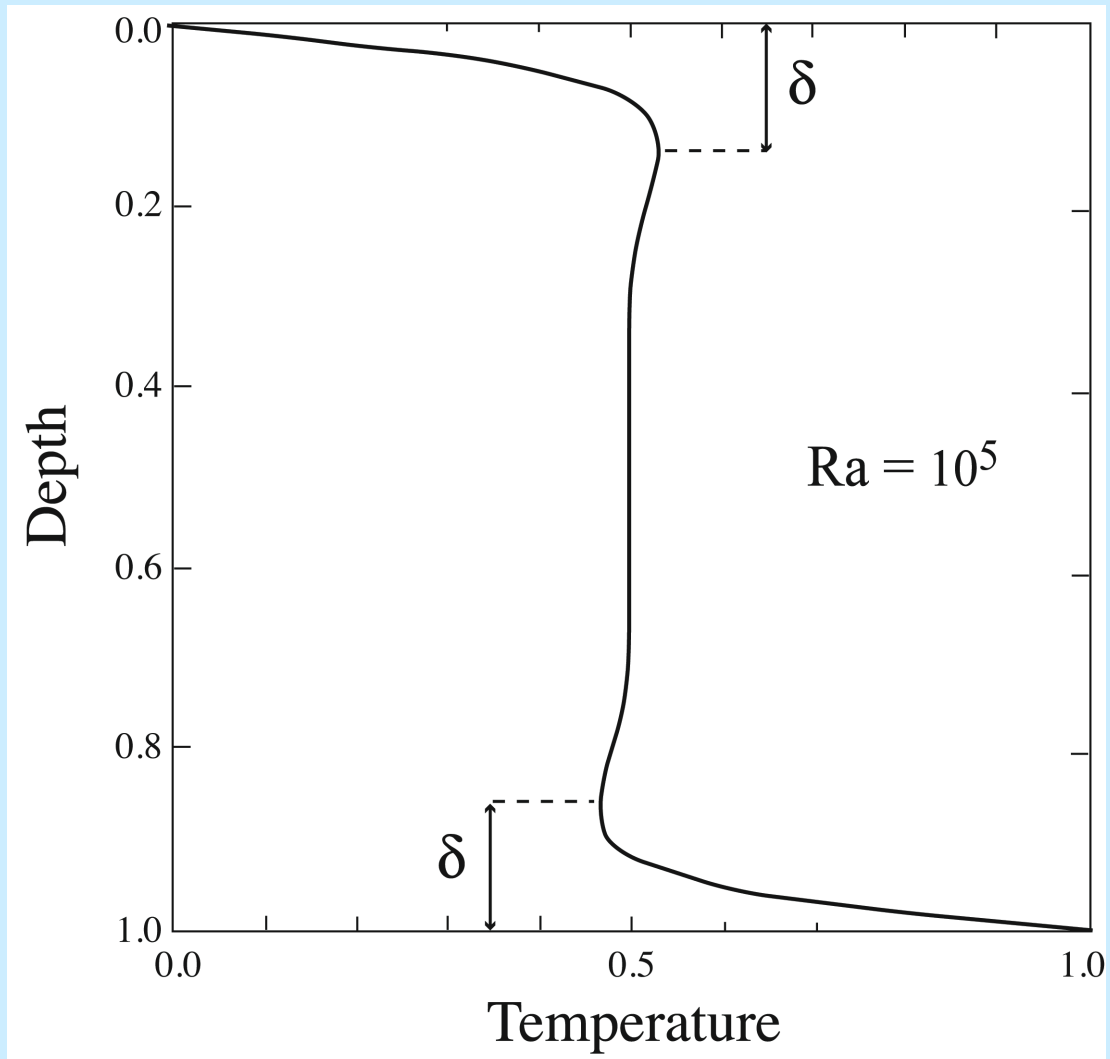
$$T = \bar{T}(z, t) + \theta(x, y, z, t).$$

$$\rho C_p \left[ \frac{\partial \bar{T}}{\partial t} + \frac{\partial \overline{w\theta}}{\partial z} \right] = \lambda \frac{\partial^2 \bar{T}}{\partial z^2}.$$

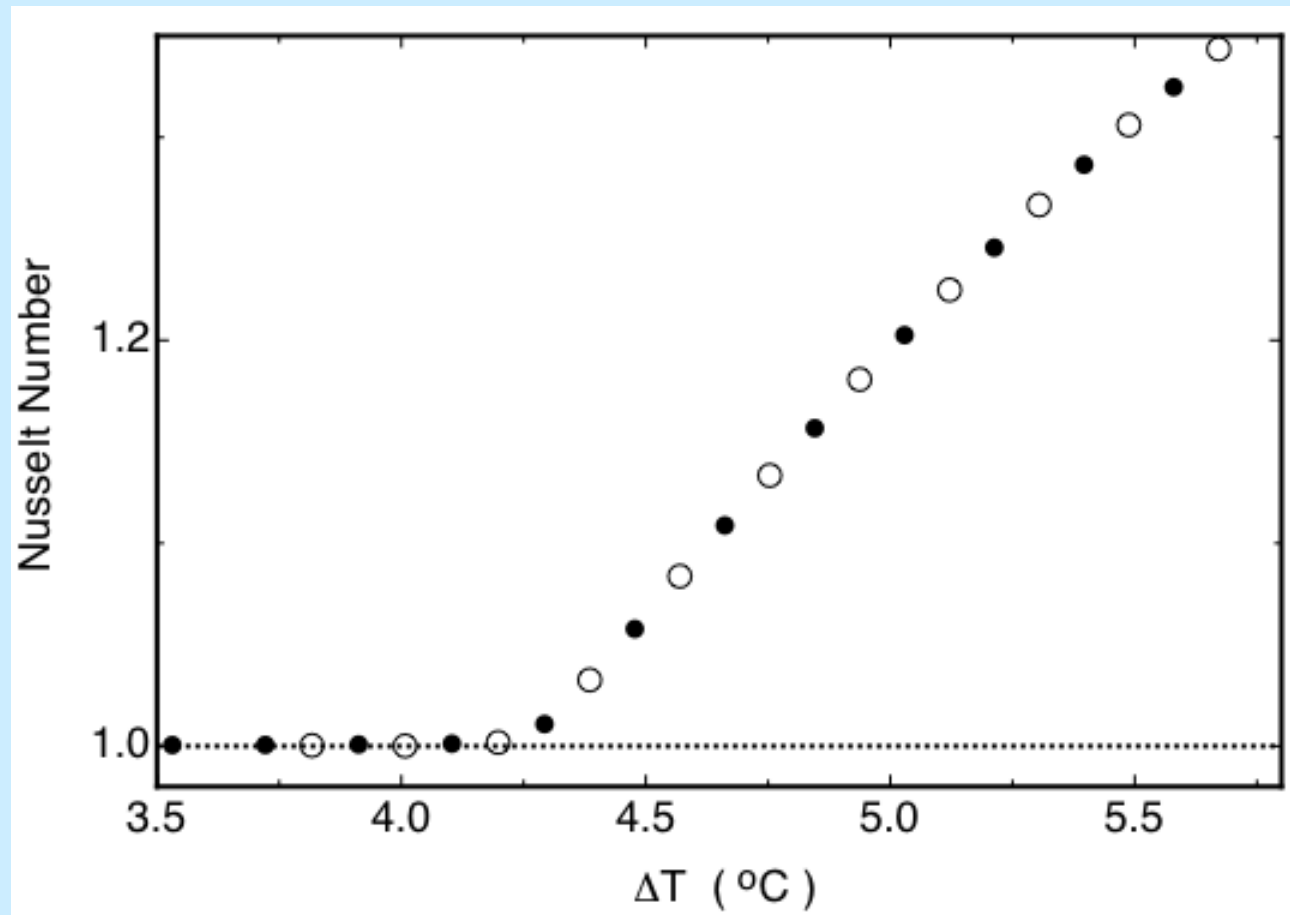
$$\rho C_p \frac{\partial \bar{T}}{\partial t} = - \frac{\partial}{\partial z} \left[ -\lambda \frac{\partial \bar{T}}{\partial z} + \rho C_p \overline{w\theta} \right]$$

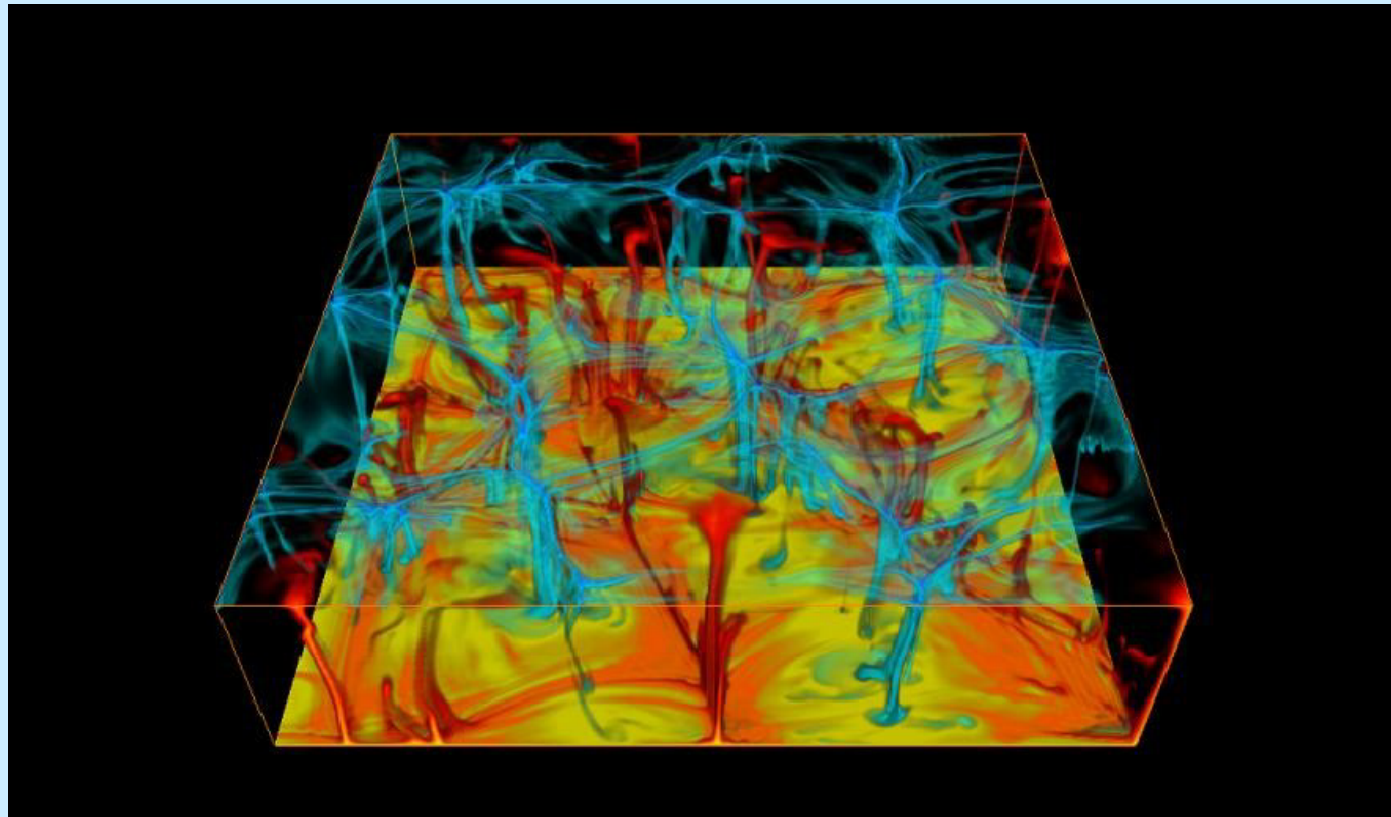
$$\rho C_p \frac{\partial \bar{T}}{\partial t} = - \frac{\partial \bar{q}}{\partial z}$$

$$\bar{q} = -\lambda \frac{\partial \bar{T}}{\partial z} + \rho C_p \overline{w\theta}.$$



$$\bar{q} = -\lambda \frac{\partial \bar{T}}{\partial z} + \rho C_p \overline{w\theta} = \text{constant} = Q,$$





$$\text{Ra} = 10^8$$

*Physical characteristics of geological convective systems*

System	$h$	$\Delta T$ , K	$\mu$ , Pa s	Pr	Ra
Upper mantle	660 km	1300 †	$5 \times 10^{20}$	$10^{23}$	$10^6$
Whole mantle	3000 km	3300 †	$5 \times 10^{21}$	$10^{24}$	$10^7$
Basaltic lava lake	50 m	50 ‡	10	$10^3$	$10^{12}$
Basaltic magma reservoir	1 km	50 ‡	10	$10^3$	$10^{16}$
Dacitic magma reservoir	1 km	50 ‡	$10^6$	$10^8$	$10^{11}$

Values have been rounded off for clarity.

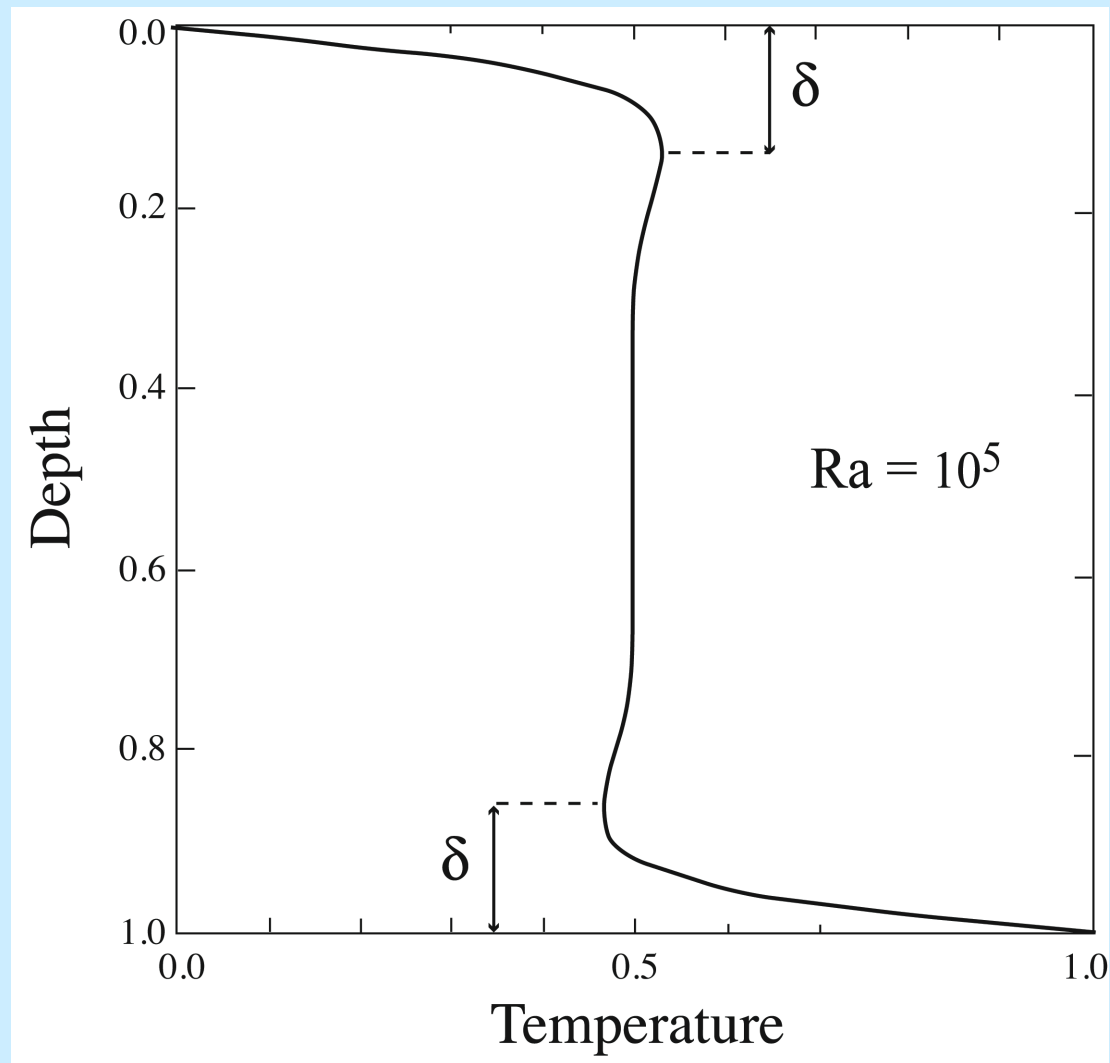
† True temperature difference deduced from the mantle geotherm of Figure 2.4.

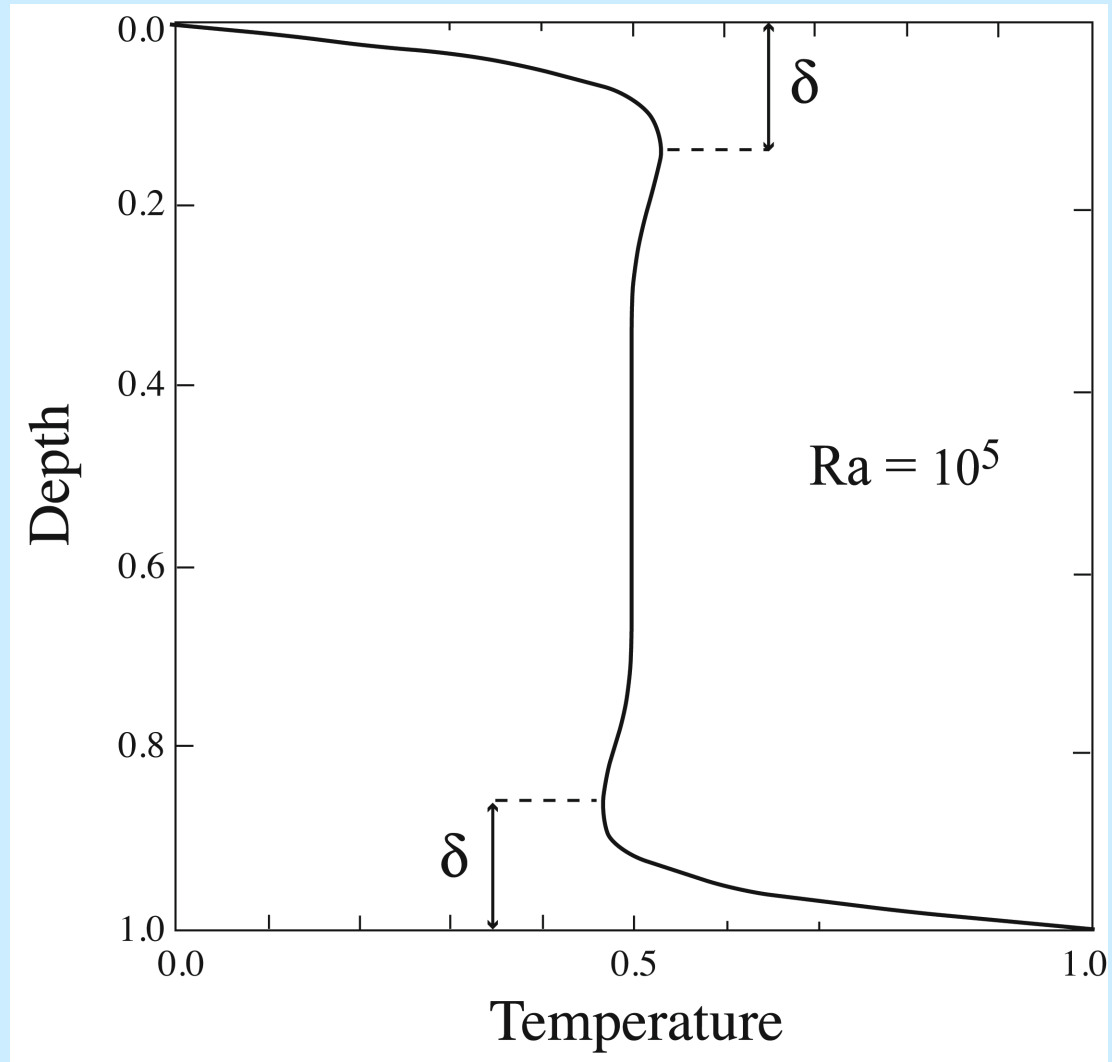
‡ temperature difference across the actively convecting part of the system.

*The “4/3” law for the convective heat flux at high Rayleigh number*

$$\text{Nu} = C_N(\text{Pr})f_N(\text{Ra}),$$







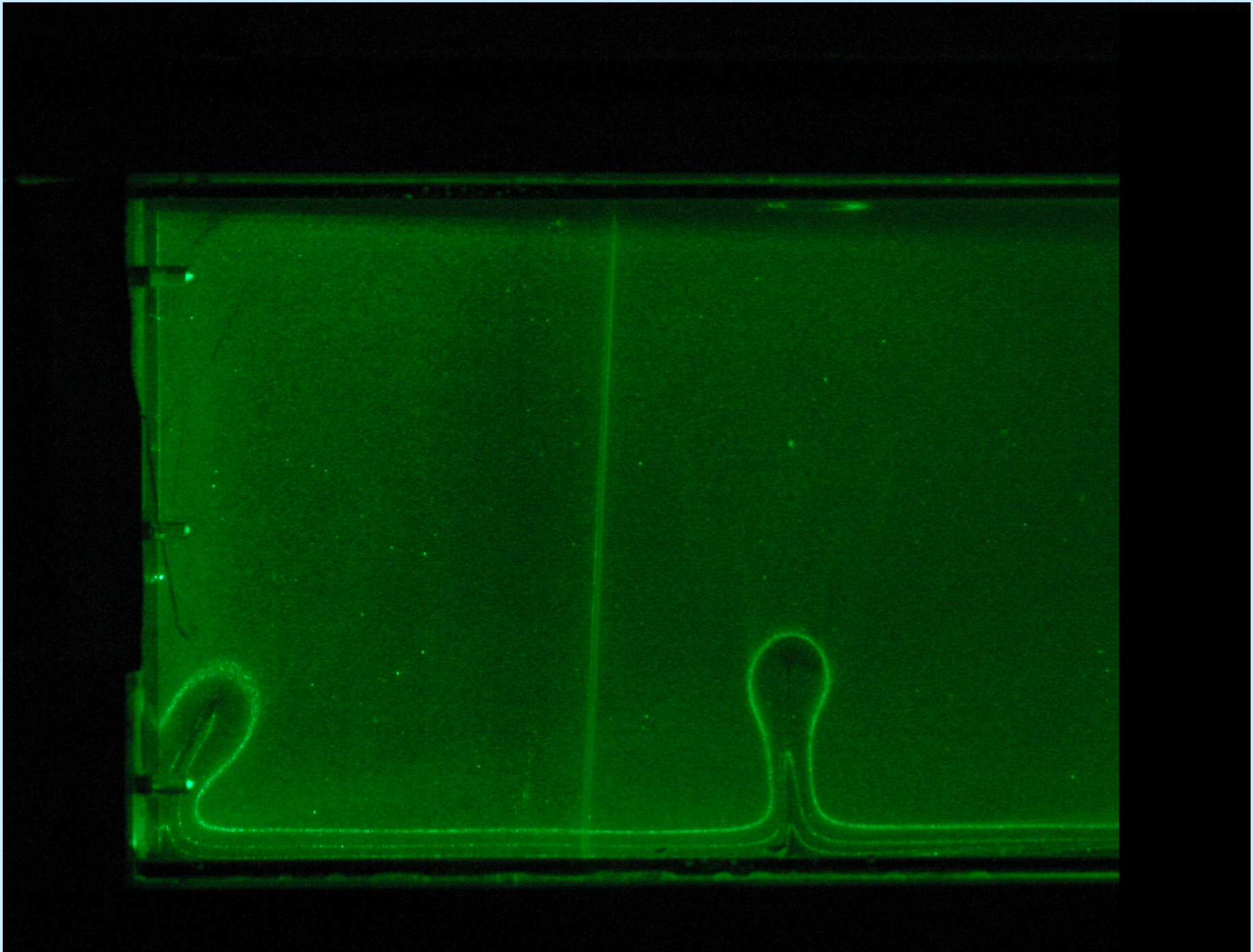
$$Q = k \frac{\Delta T}{2\delta},$$

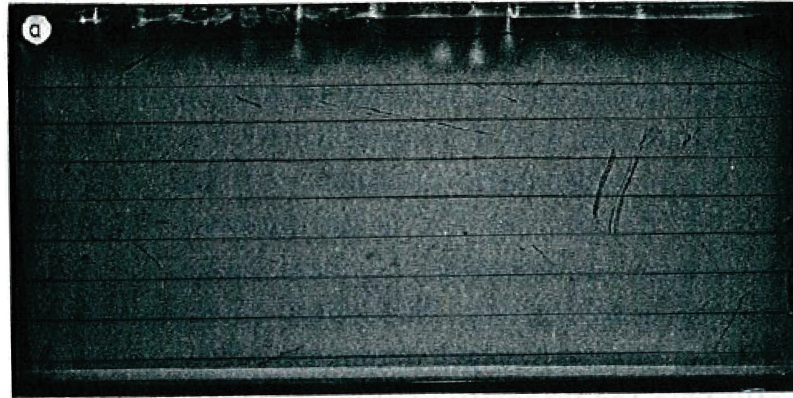
*The “4/3” law for the convective heat flux at high Rayleigh number*

$$\text{Nu} = C_N(\text{Pr})f_N(\text{Ra}),$$

$$Q = k \frac{\Delta T}{2\delta},$$

$$\text{Nu} = \frac{h}{2\delta}.$$





$$Q = \text{Nu} \cdot k \frac{\Delta T}{h} = C_N (\text{Pr}) k \frac{\Delta T}{h} f_N \left( \frac{\rho_0 g \alpha \Delta T h^3}{\kappa \mu} \right),$$

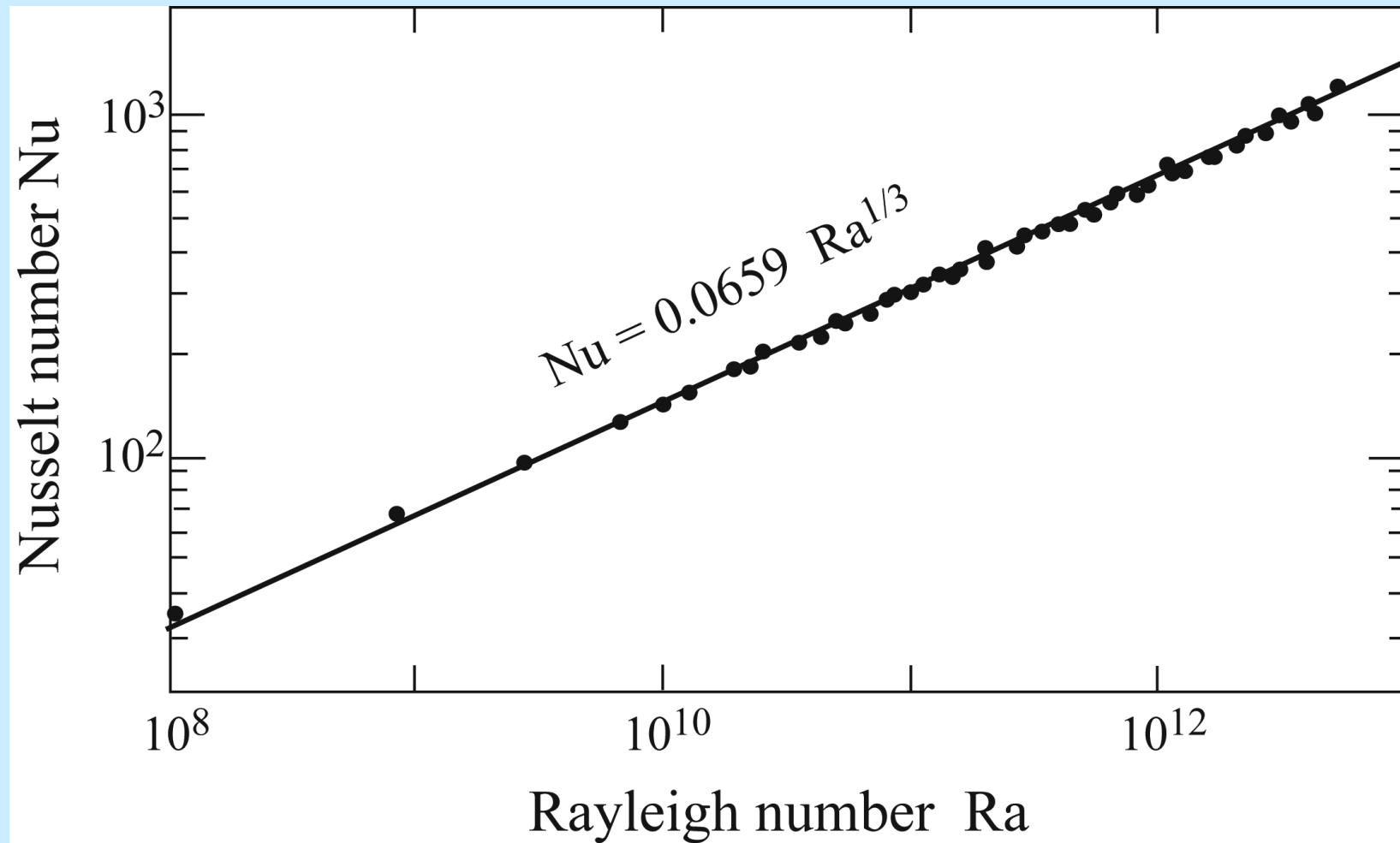
$$f_N(\text{Ra}) \propto \text{Ra}^{1/3}$$

$$\text{Nu} = C_N \text{Ra}^{1/3}.$$

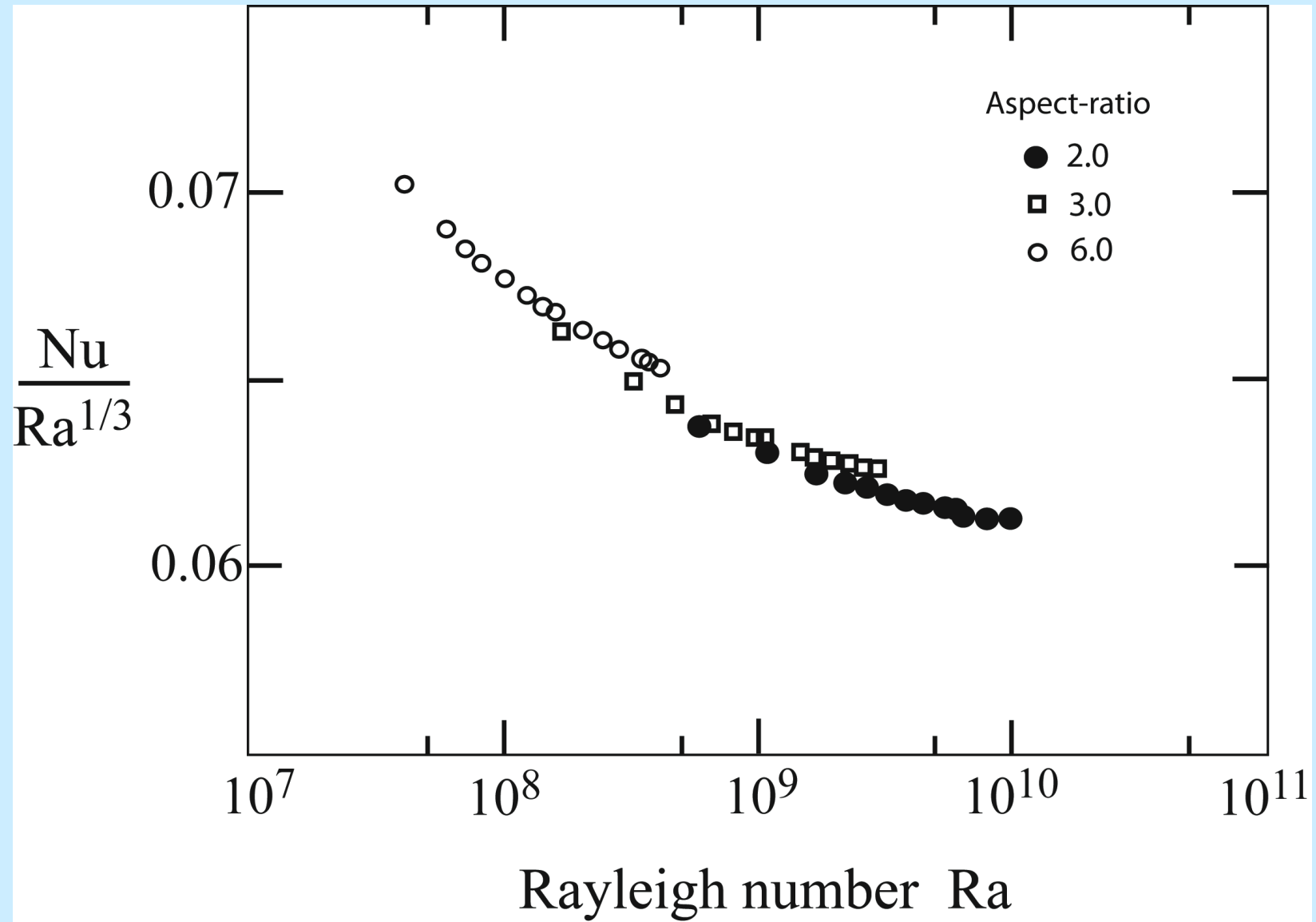
$$Q = C_Q k \left( \frac{g \alpha}{\kappa \nu} \right)^{1/3} \Delta T_\delta^{4/3},$$

$$C_Q = 2^{4/3} C_N$$

Rigid boundaries,  $Pr = 2750$ ,  $L/h > 1$  (Goldstein et al., 1990)



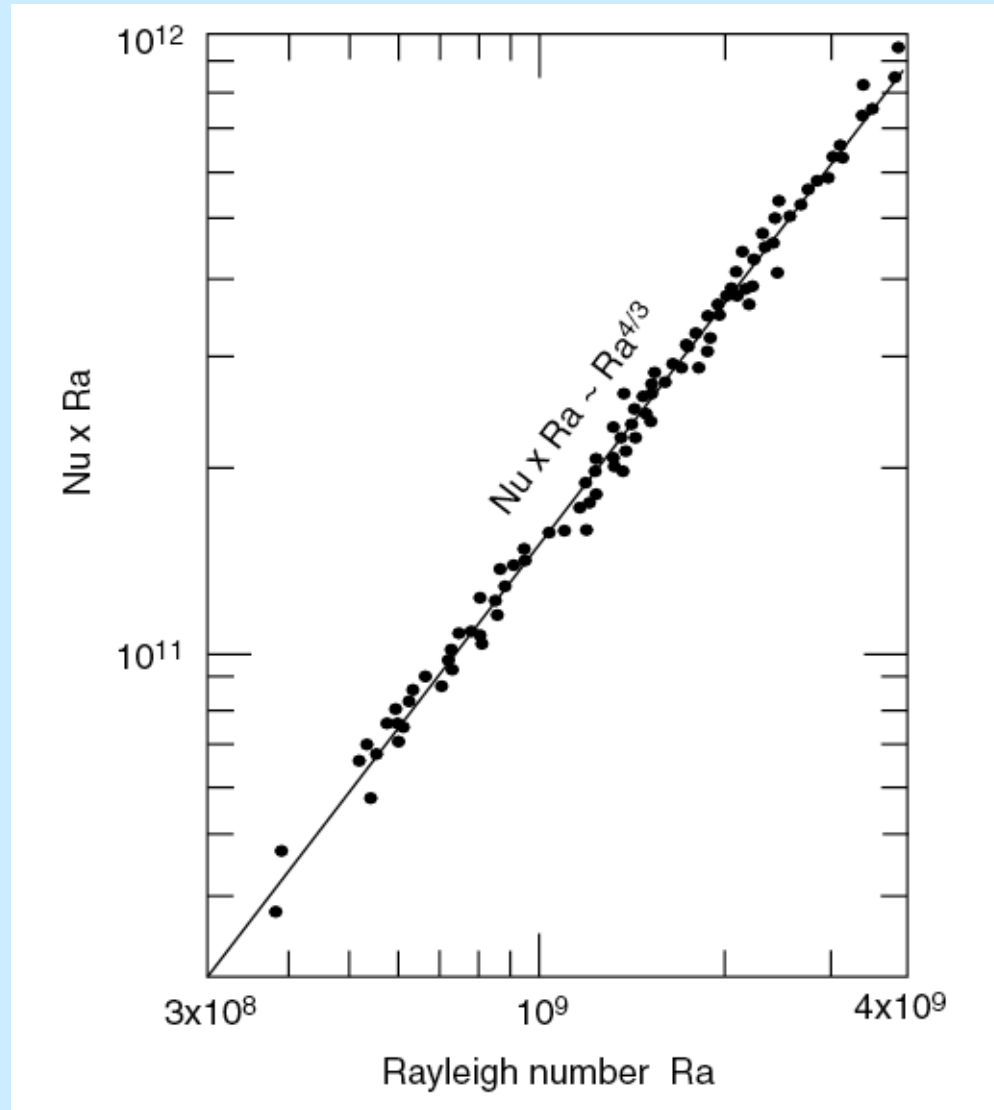
Rigid boundaries,  $Pr \approx 5$  (water) (Funfschilling et al., 2005)





Free upper boundary (cooling from top only),  $Pr \approx 5$  (water)

$L/h = 1$  {Katsaros et al., 1977}



*. Data on the convective heat flux in Rayleigh–Benard convection*

Pr	Ra	Bound. cond.	Aspect ratio	$C_N$ <sup>+</sup>	$C_Q$ <sup>++</sup>	Reference
4–6	$5 \times 10^{10} - 10^{11}$	Rigid	0.98	0.060	0.15	(Funfschilling <i>et al.</i> , 2005)
4–6	$10^{10}$	Rigid	2	0.062	0.16	(Funfschilling <i>et al.</i> , 2005)
4–6	$3 \times 10^8 - 4 \times 10^9$	Free	1	§	0.16†	(Katsaros <i>et al.</i> , 1977)
2750	$10^8 - 10^{13}$	Rigid	> 1	0.0659	0.17	(Goldstein <i>et al.</i> , 1990)
$\infty$	$10^6 - 10^9$ ‡	Free	1.5	0.150¶¶	0.378¶¶	(Hansen <i>et al.</i> , 1992)

<sup>+</sup> Constant in the Nu versus Ra relationship (5.81).

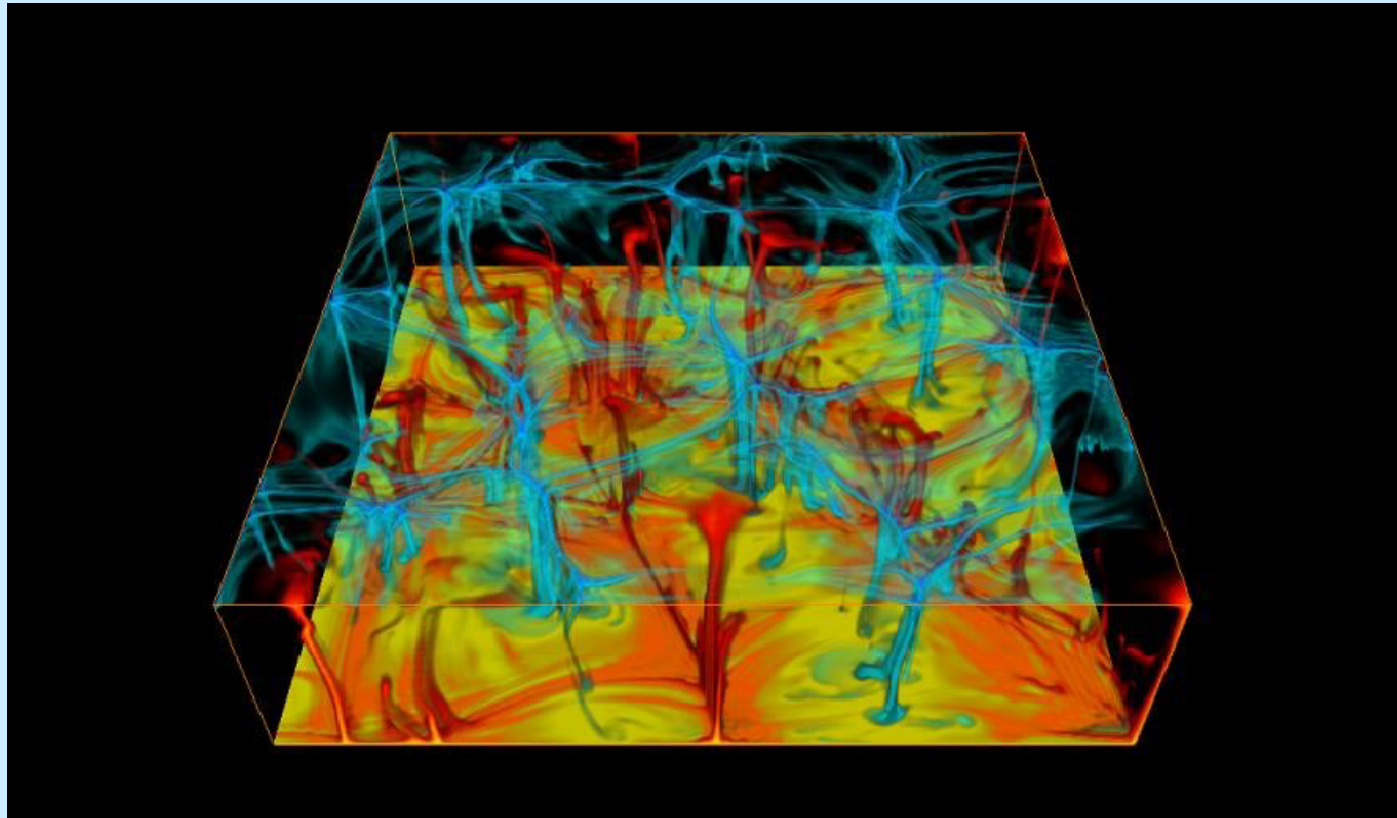
<sup>++</sup> Constant in the local heat flux scaling law (5.82).

§ Only the boundary layer scaling can be determined in this transient cooling experiment.

† Value re-calculated for a 1/3 scaling exponent (instead of 0.33).

¶¶ Value re-calculated for a 1/3 scaling exponent at  $Ra = 10^9$ .

‡ Numerical calculations in 2D.



$$\mathbf{Ra} = 10^8$$

$$0 = -\nabla P_h + \rho_o [1 - \alpha(\bar{T} - T_o)] \mathbf{g},$$

$$\rho_o \mathbf{v} \cdot \frac{\partial \mathbf{v}}{\partial t} + \rho_o \mathbf{v} \cdot (\mathbf{v} \nabla \mathbf{v}) = -\mathbf{v} \cdot \nabla P - \mathbf{v} \cdot (\nabla \cdot \boldsymbol{\tau}) + \alpha \rho_o w \theta g,$$

$$e_c = v^2 / 2$$

$$\begin{aligned} \rho_o \frac{\partial e_c}{\partial t} + \nabla \cdot (\rho e_c \mathbf{v}) = & -\nabla \cdot (\mathbf{v} p) + p \nabla \cdot \mathbf{v} \\ & - [\mathbf{v} \cdot (\nabla \cdot \boldsymbol{\tau}) - \nabla \cdot (\boldsymbol{\tau} \cdot \mathbf{v})] - \nabla \cdot (\boldsymbol{\tau} \cdot \mathbf{v}) + \alpha \rho_o w \theta g, \end{aligned}$$

## Irreversible kinetic dissipation

$$\psi = \mathbf{v} \cdot (\nabla \cdot \boldsymbol{\tau}) - \nabla \cdot (\boldsymbol{\tau} \cdot \mathbf{v}).$$

$$\int_V \rho_o \frac{\partial e_c}{\partial t} dV + \int_S \rho_o e_c \mathbf{v} \cdot dS = - \int_S p \mathbf{v} \cdot dS + \int_V p \nabla \cdot \mathbf{v} dV$$
$$- \int_V \psi dV - \int_S (\boldsymbol{\tau} \cdot \mathbf{v}) \cdot dS + \int_V \alpha \rho_o w \theta g dV.$$

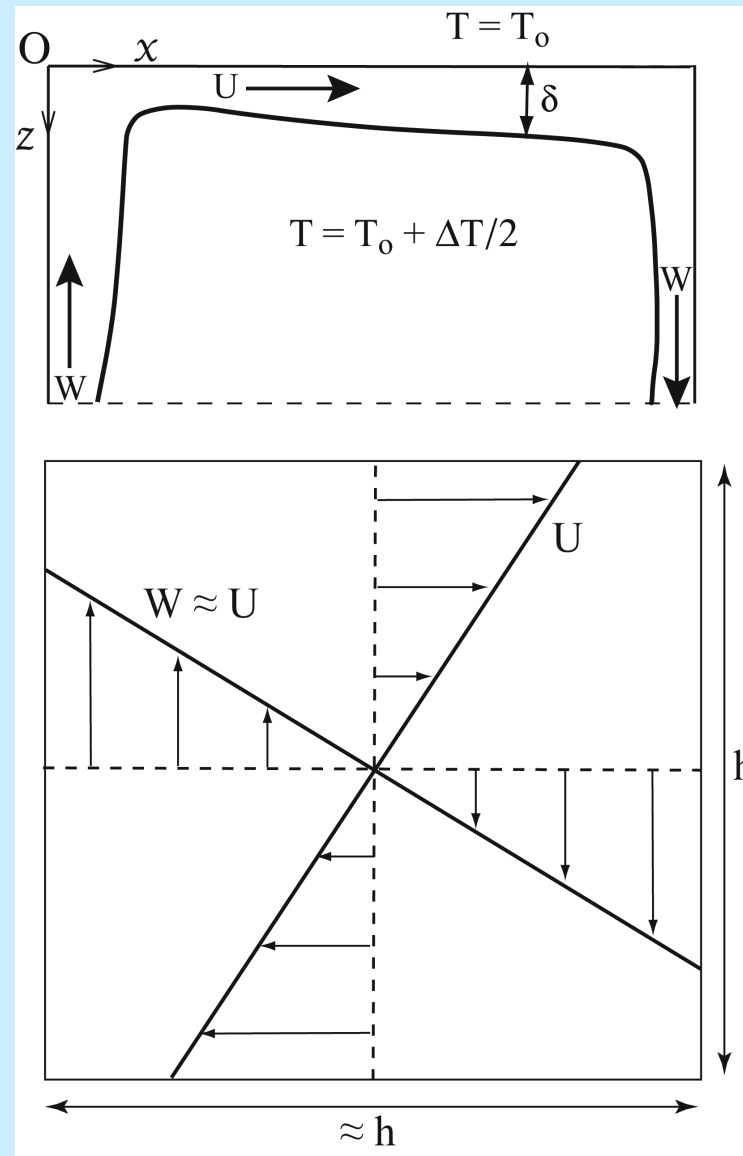
$$+ \int_V \rho \alpha g w \theta dV - \int_V \psi dV = 0.$$

$$+ \int_0^h \rho \alpha g \bar{w} \bar{\theta} dz - \int_0^h \bar{\psi} dz = 0.$$

$$\begin{aligned} \int_0^h \rho C_p \bar{w} \bar{\theta} dz &= \int_0^h \left( Q + k \frac{d\bar{T}}{dz} \right) dz \\ &= Qh + k [T(h) - T(0)] = Qh - k \Delta T. \end{aligned}$$

$$\epsilon = \int_0^h \bar{\psi} dz = \mu \frac{\nu^2}{h^3} (\text{Nu} - 1) \text{RaPr}^{-2}.$$

# Free boundaries, large Prandtl number



$$\int_0^h \psi dz \sim \mu \frac{U^2}{h^2} h \sim \mu \frac{v^2}{h^3} \text{Re}^2.$$

$$\mu \frac{v^2}{h^3} (\text{Nu} - 1) \text{RaPr}^{-2} \sim \mu \frac{v^2}{h^3} \text{Re}^2,$$

Remember that  $Nu \sim h/\delta$ .

Add balance between horizontal advection  
and vertical diffusion  
(remember scalings for laminar plumes)

$$\delta \sim \sqrt{\kappa h / \dot{U}}.$$

$$\text{Re}^{1/2} \text{Pr}^{1/2} \sim \text{Nu}.$$



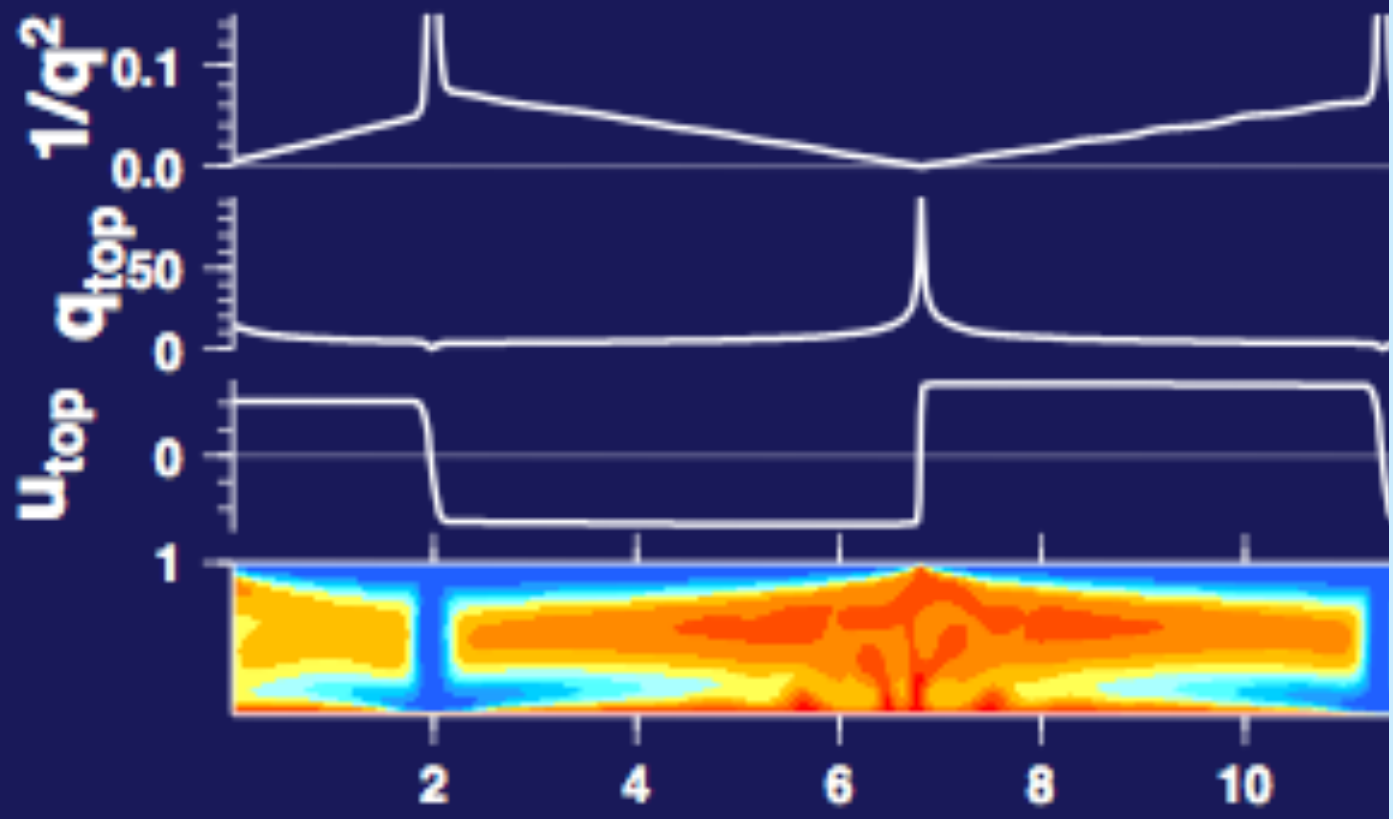
$$\text{Nu} \sim \text{Ra}^{1/3}, \quad \text{Re} \sim \text{Ra}^{2/3} \text{Pr}^{-1},$$

$$\delta \sim h \text{Ra}^{-1/3}, \quad U \sim \frac{\kappa}{h} \text{Ra}^{2/3}.$$

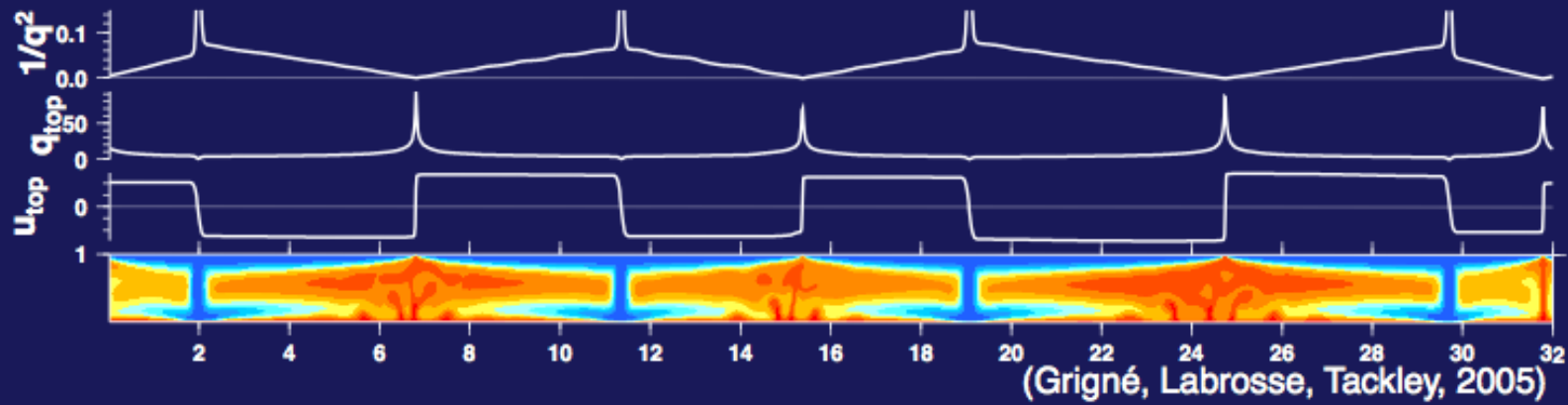
$$\delta \sim \sqrt{\kappa h / \dot{U}}$$

This is the average boundary layer thickness. The local thickness at distance  $x$  from upwelling is

$$\delta \sim \sqrt{\kappa x / \dot{U}}$$



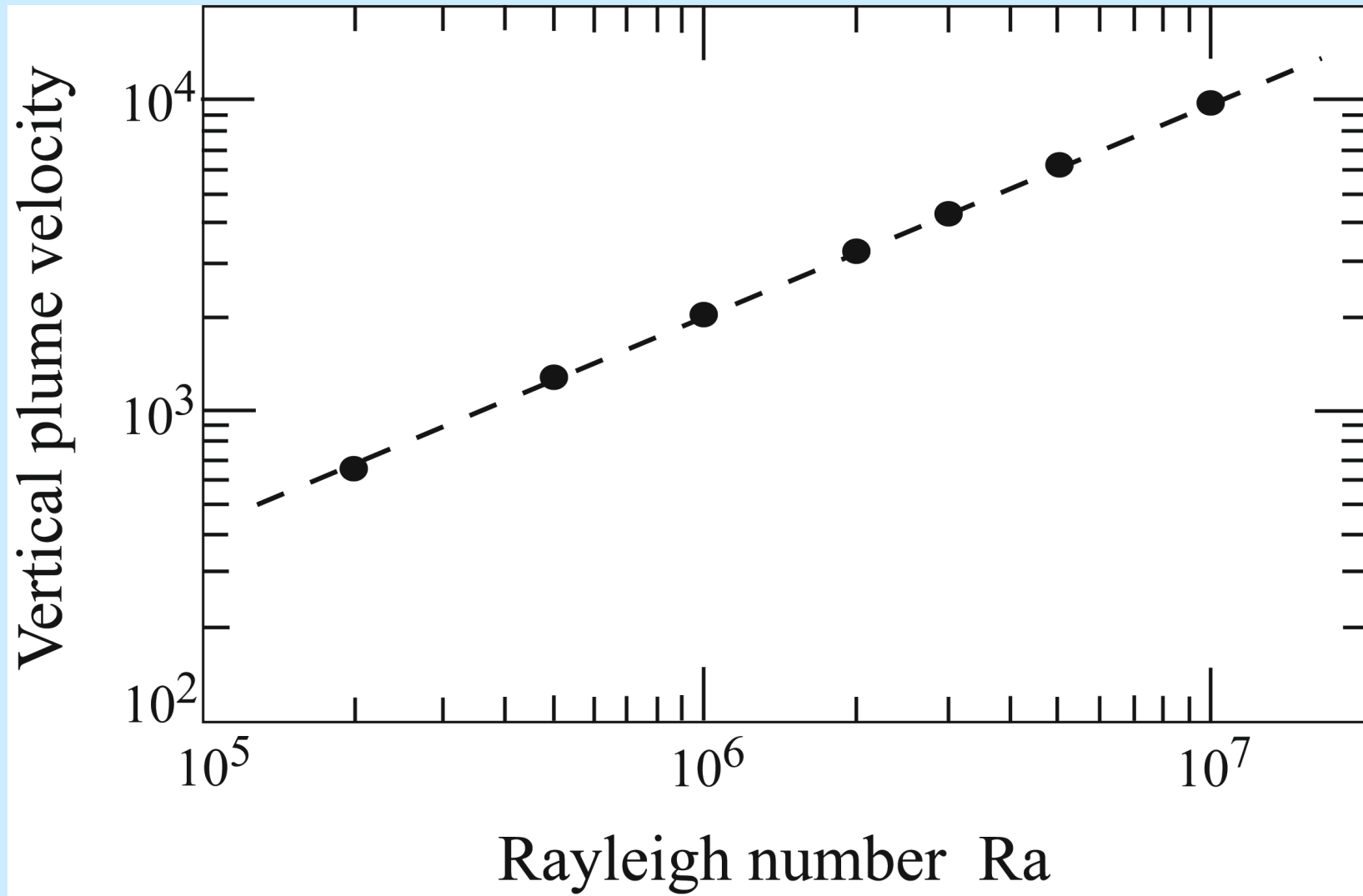
# Plate tectonics in 2D



$$\text{Heat flux} \sim (\text{distance})^{-1/2}$$

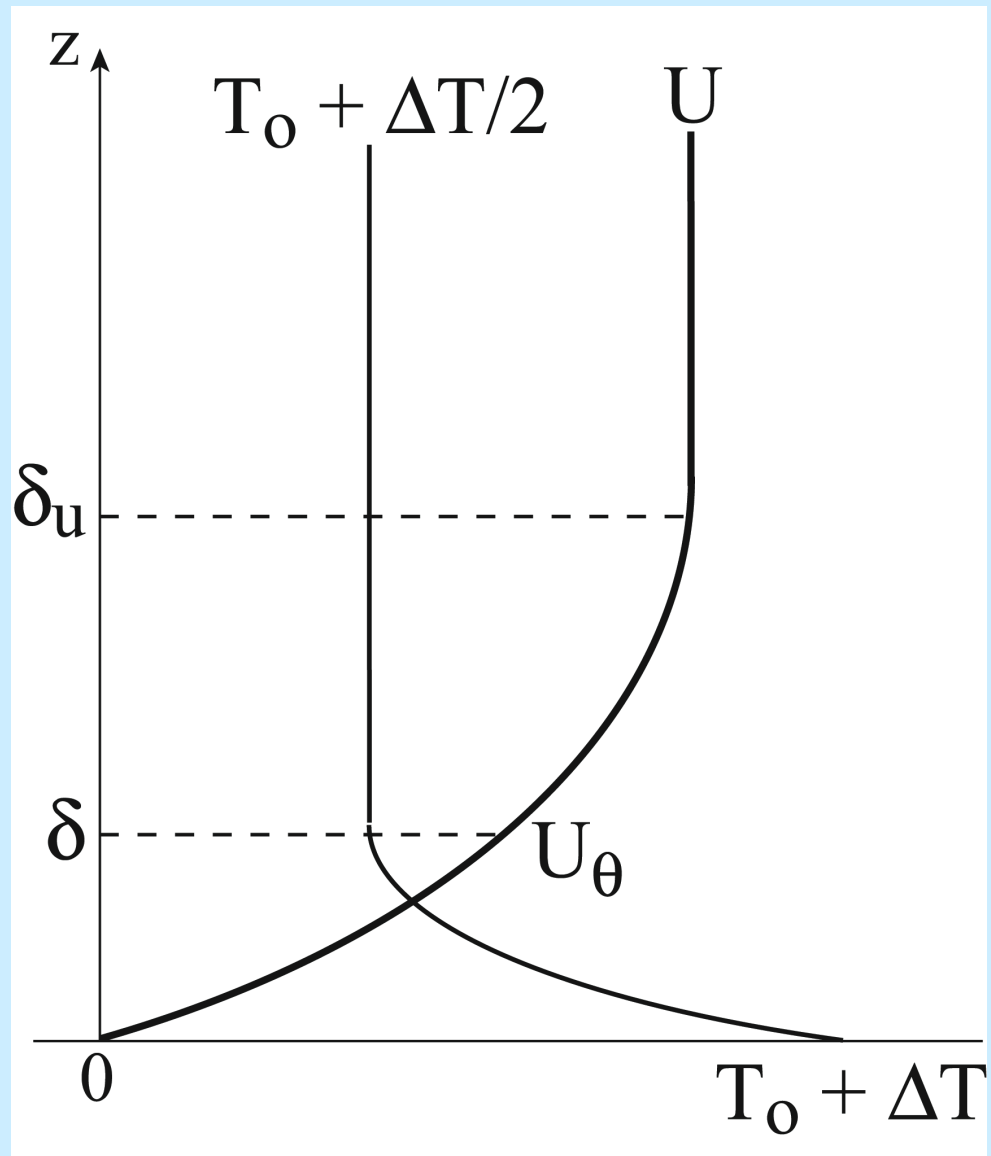
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(Galsa & Lenkey, Phys. Fluids 2007)

With rigid boundaries



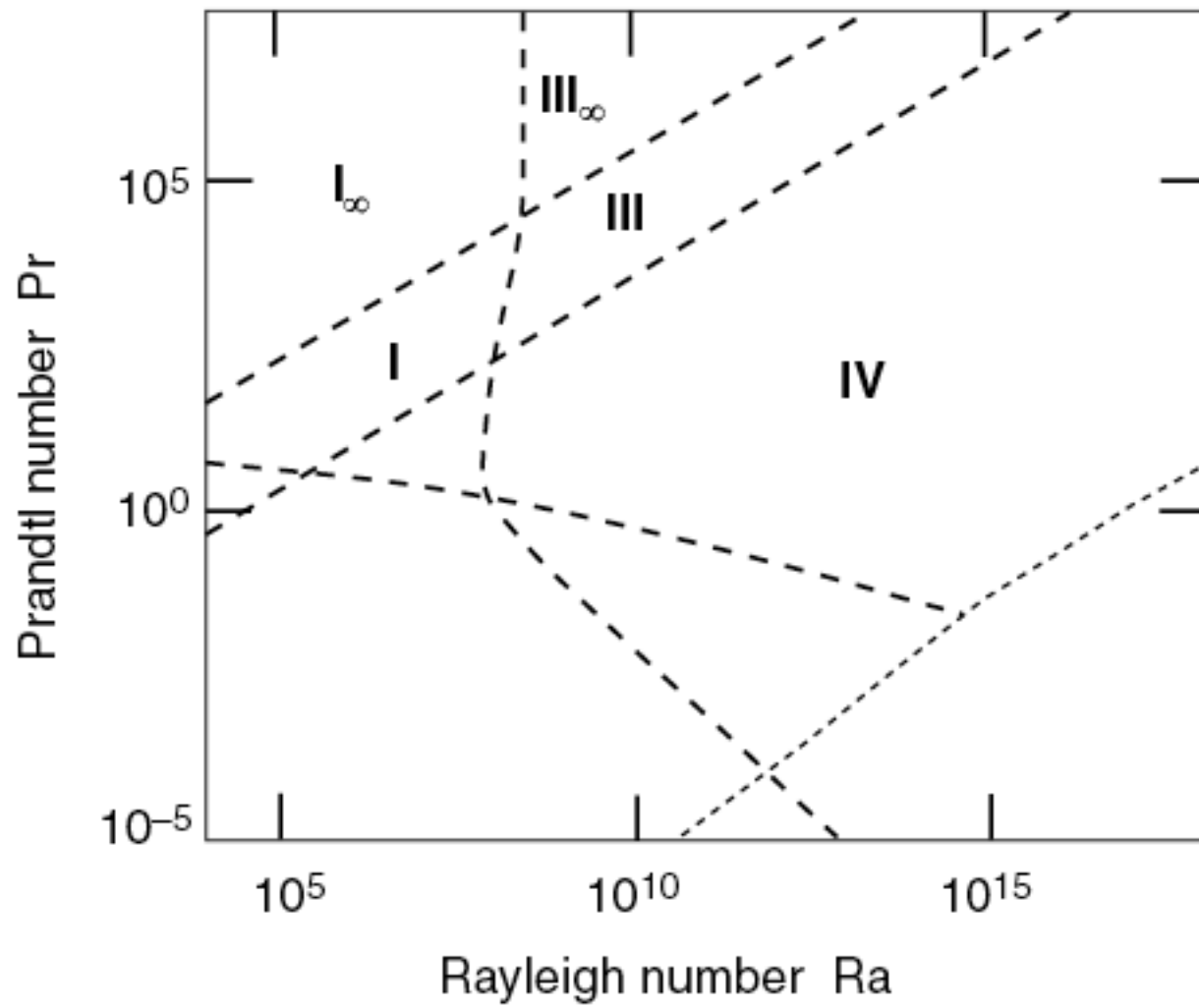
# The regimes of Rayleigh-Bénard convection in rigid enclosures

(Grossmann & Lohse, J. Fluid Mech. 2000, Phys. Fluids. 2001)

Regime	Dominant dissipation §	Nu	Re
I	$(u, B) - (\theta, B)$	$0.31 \text{ Ra}^{1/4} \text{ Pr}^{-1/12}$	$0.073 \text{ Ra}^{1/3} \text{ Pr}^{-5/6}$
$I_\infty$ ( $\text{Pr} \gg 1$ )	$(u, B) - (\theta, B)$	$0.35 \text{ Ra}^{1/5}$	$0.054 \text{ Ra}^{3/5} \text{ Pr}^{-1}$
III	$(u, B) - (\theta, I)$	$0.018 \text{ Ra}^{3/7} \text{ Pr}^{-1/7}$	$0.023 \text{ Ra}^{4/7} \text{ Pr}^{-6/7}$
$III_\infty$ ( $\text{Pr} \gg 1$ )	$(u, B) - (\theta, I)$	$0.027 \text{ Ra}^{1/3}$	$0.015 \text{ Ra}^{2/3} \text{ Pr}^{-1}$
IV	$(u, I) - (\theta, I)$	$0.060 \text{ Ra}^{1/3}$	$0.088 \text{ Ra}^{4/9} \text{ Pr}^{-2/3}$

§Dominant contributions to kinetic and thermal dissipation (see text).  $u$  and  $\theta$  stand for the kinetic and thermal dissipation, respectively, and symbols  $I$  and  $B$  indicate interior and boundary-layer contributions, respectively.





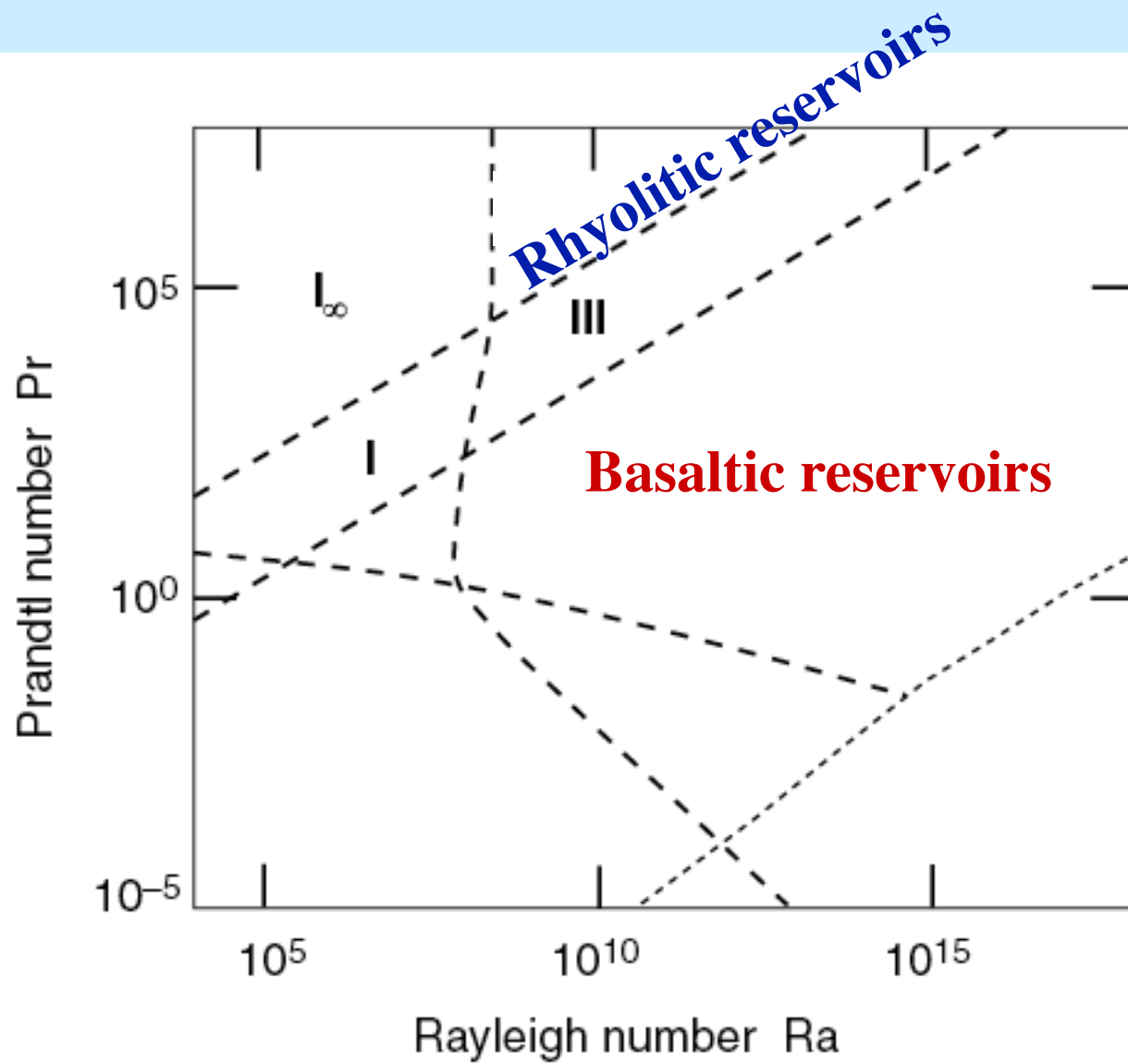
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