





2240-20

Advanced School on Scaling Laws in Geophysics: Mechanical and **Thermal Processes in Geodynamics**

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DEEP CONVECTION (Compressible Mantle Convection)

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Compressible mantle convection



1) Adomy-Williamson relation and Pressure P and entropy

5, PCS, P).

$$\nabla P(s,p) = \left(\frac{\partial Pr}{\partial s}\right) p + \left(\frac{\partial Pr}{\partial p}\right)_{s} \nabla P. \qquad 0$$

Assume an adiabatic process,
$$\nabla S = 0$$
. $\nabla P = -P_r g \stackrel{!}{=} r$

$$\frac{1}{P_r} \frac{dP_r}{dr} = -P_r \left(\frac{\partial P_r}{\partial P}\right)_S$$

Ks = Pr (2P) 5 dofines the isentropec balk moduli

$$\frac{1}{Pr}\frac{dPr}{dr} = -\frac{Pr8}{ks}$$

Define Gruneisen's parameter $T = \frac{\alpha K_s}{PC_p}$, and

(3) becomes
$$\frac{1}{Pr} \frac{dPr}{dr} = -\frac{g\alpha}{\Gamma C_p}$$
 (4)

Assume that parameters on the RHS of @ are constant, and integrate @.

$$P_r(r) = P_o e^{\frac{g\alpha}{Tc_p}(d-r)}$$

where po is the surface density at r=d.

Reference density or uncreases with pressure or depth, 7.e., the fluid is compressible. This affects manufle convection.

2) Governing equation for compressible convections D.(PV)=0

-∇P+D.T-8P8 e3=0

D

X3

LB

×1 PrGp(Jt + V. VT) + PrBXTV3 = V(KVT) + TEJUSJ +PrH

where $\delta \rho = \rho_r \left(-\alpha (T - T_r) + k_T^{-1} \rho \right)$ and p is the dynamic pressure and ky is isothered modulus, Tr is reference temperature.

Non-domension live the agustions, uses d, to, AT, Po, Mo, Qo, Cp-o, To, and to to normalize x, v, T, P, n, x, cp, T and k. Fine, pressure and internal heaty scales an $\frac{d^2}{x_o}$, $\frac{\eta_o \kappa_o}{d^2}$, and $\frac{\kappa_o \Delta T}{d^2}$, respectively. D.(PrV)=0
3
-∇ρ+ ∇. + [Prαg Ra(T-Tr) - αβρ; σ=0 (0)

PrCp(2T+V.VT)+prgaDilT+Ts)U3 = V.(KVT) + Di T. E + PrH (1)

$$D_{\tau} = \frac{\alpha_0 f_0 d}{c_{p,o}}$$
 Dissipatan number

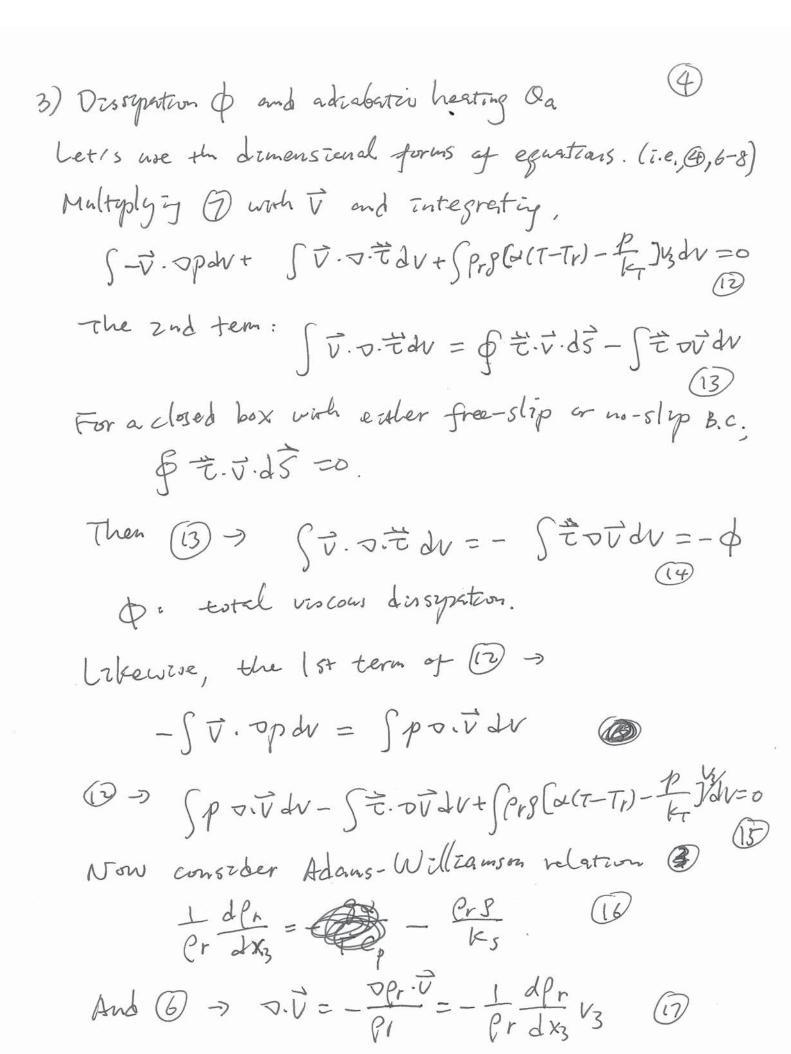
 $Ra = \frac{C_0 f_0 \alpha_0}{\kappa_0 \gamma_0}$

Notice that if y we assume constant parameters such that gravitational acceleration are constant, $\alpha = \beta = c_p = T = k = 1$ in the non-dimensional equations (3), (0) and (1).

Remarks:

A)
$$T_{ij} = \eta \left(\frac{\partial V_i}{\partial x_j} + \frac{\partial V_j}{\partial x_i} \right) - \frac{2}{3} \eta \delta_{ij} \frac{\partial V_k}{\partial x_k} + \eta_2 \delta_{ij} \frac{\partial V_k}{\partial x_k}$$
 $\eta_2 : \text{bulk viscosiny, often ignored.}$
 $T_{ij} = \eta \left(\frac{\partial V_i}{\partial x_j} + \frac{\partial V_j}{\partial x_i} \right) - \frac{2}{3} \eta \delta_{ij} \frac{\partial V_k}{\partial x_k}$

$$\frac{dT_r}{dr} = -\frac{\alpha T_r \beta}{c_p}$$
, adiabatic tapperature gradient.



total viscous dissyran = estal adiabatis heating

Remark: The effect of p on SP is the key to

assuring the balance [Leng and Thong 2008]

From (i), $\phi = \int \frac{g_{\alpha}d}{dc_{p}} \cdot C_{p} c_{p} T v_{3} dv$ $\phi = \int \frac{g_{\alpha}d}{dc_{p}} \cdot C_{p} c_{p} T v_{3} dv$ $\phi = D_{2} \int \int C_{p} c_{p} c_{p} T v_{3} dv$ $\phi = D_{2} \int \int \int C_{p} c_{p} c_{p} T v_{3} dA dx_{3}$ $\phi = D_{2} \int \int \int \int C_{p} c_{p} c_{p} T v_{3} dA dx_{3}$

SA (r Cp TV3 dA is convertise heat flux and in the convection-dominated reprons (i.e., outside the type and bottom thermal boundaries) it is equal to F5 that is uniform at all depths in the convection-dominated regions and is equal to the surface Next flux (H=0),

 $2i\partial \rightarrow \Phi = Di \int_{d} F_{5}(d-28)$ $\frac{\Phi}{F_{5}} = Di \left(1-2\frac{\delta}{d}\right) \approx D_{7} \quad \text{for } 8 \ll d, \text{ or } 2 \text{ agge } Rq.$ $\frac{\Phi}{F_{5}} \approx D_{7} \qquad 2 \text{ agge } Rq.$

Hewith et al. C1975) proposed that \$=Di(1-23)
3: internel heaty rotio.