

# Two examples of use of the hierarchy of models

Brian Hoskins

Director, Grantham Institute for Climate Change, Imperial College London  
Professor of Meteorology, University of Reading

# The Hierarchy of Models

## Use

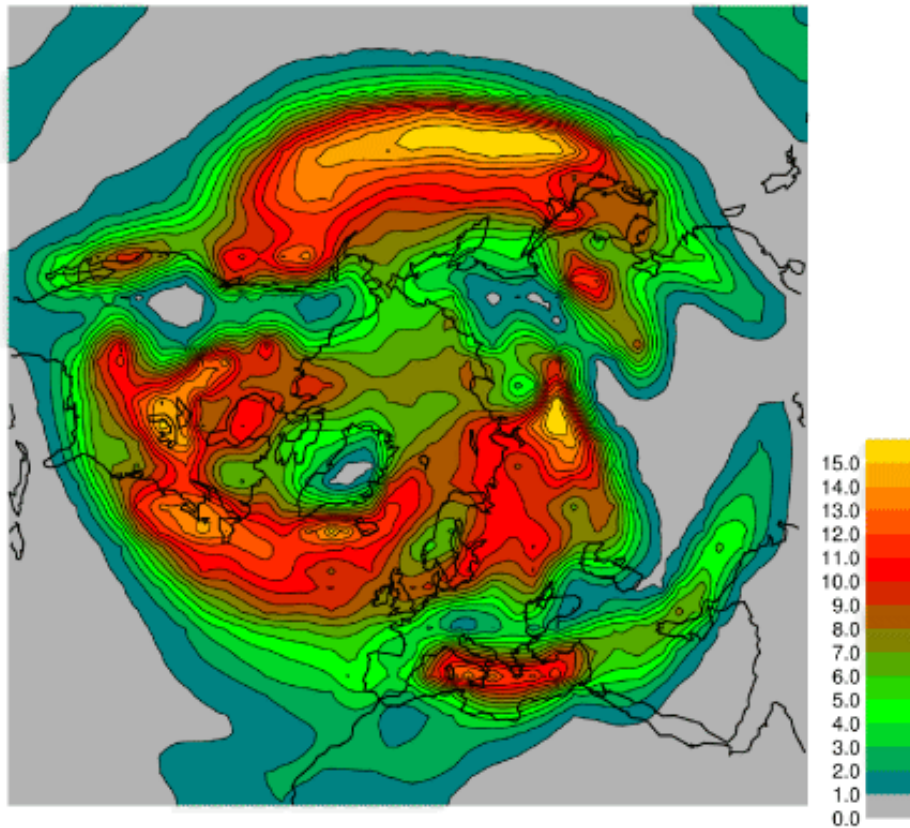
- To aid in the diagnosis of more complex models
- To aid in determination of physical significance
- To provide a framework for analysing observational data
- To increase the theoretical base and build understanding
- To aid the development of more complex models

## Members of the Hierarchy

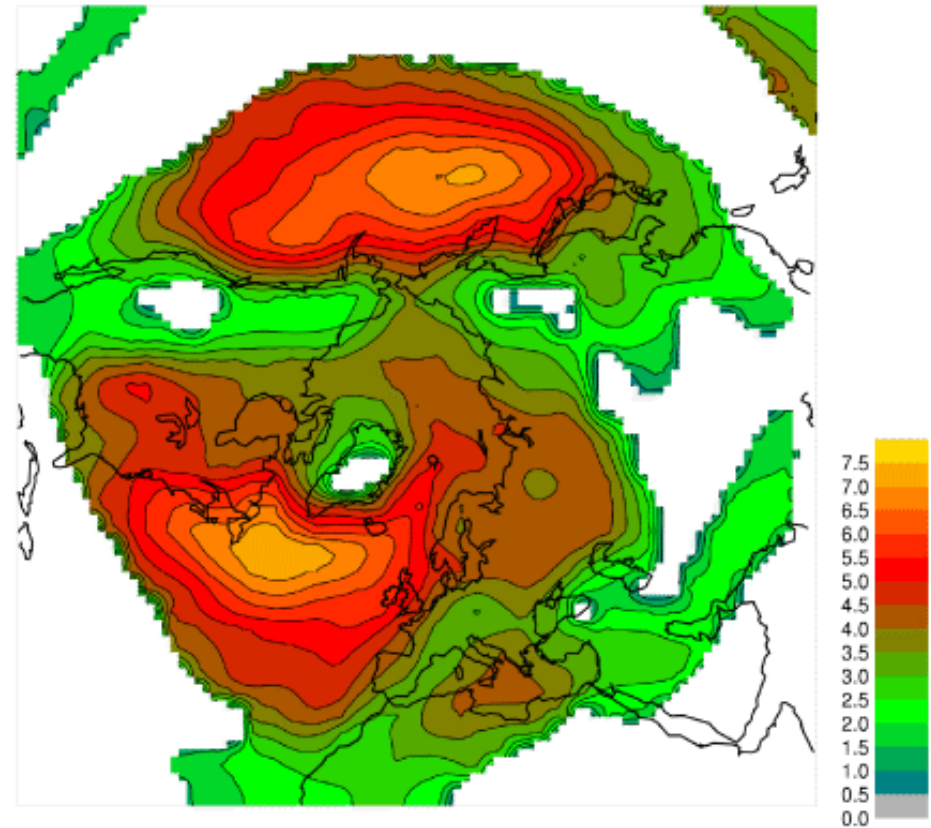
- Reduced or simplified representation of physical  
(chemical or biological) processes
- Simplified geometry
- Simplified dynamics
- Simplified domain
- Focus on a particular process or problem

# $\xi_{850}$ Tracking Statistics for DJF

Hoskins & Hodges (2002)



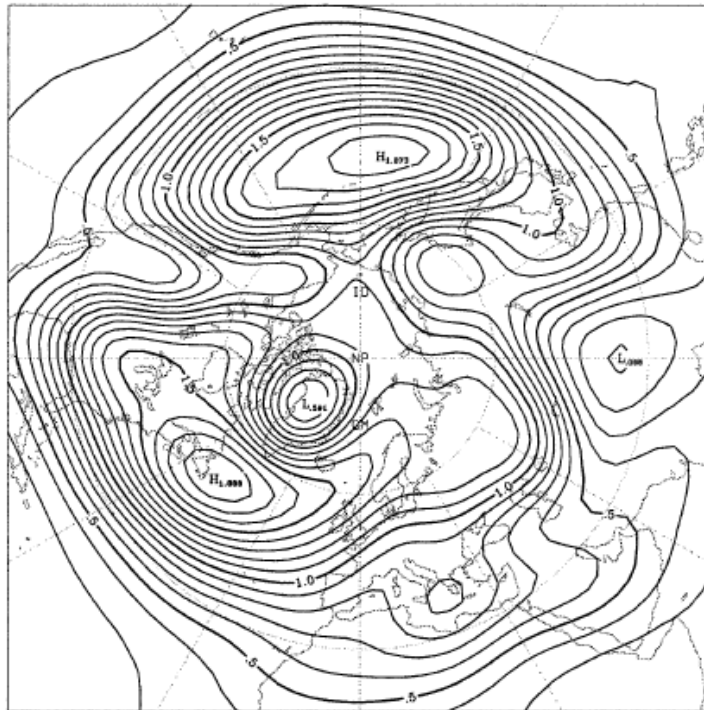
*Track Density*



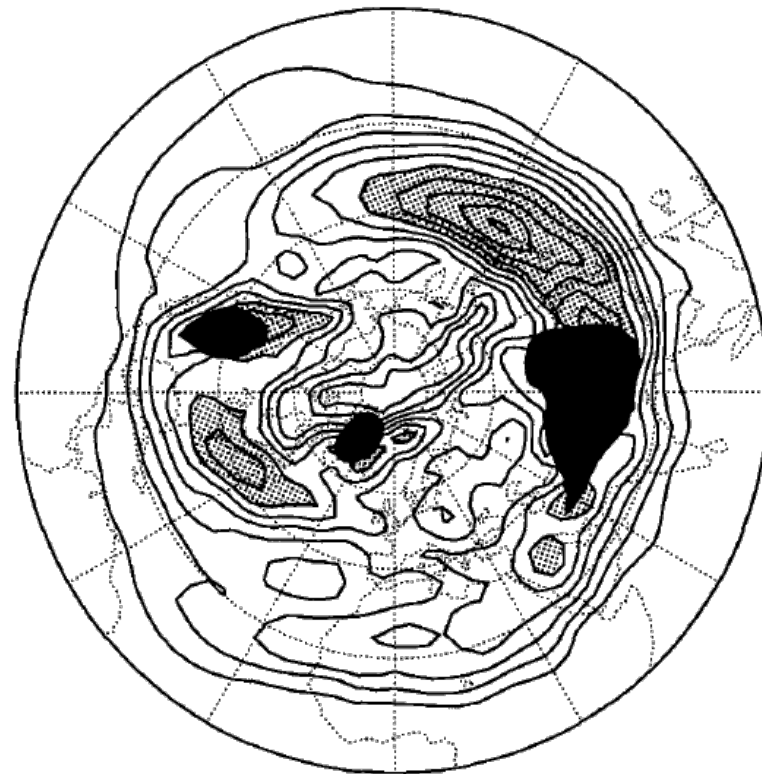
*Mean Intensity*

# Why do storm-tracks exist?

*Hoskins and Valdes, 1990*



$\xi'^2_{850}$  2-6 day

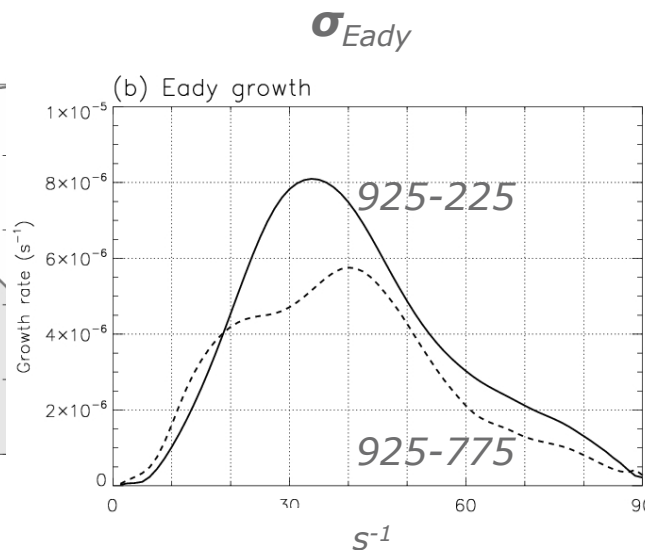
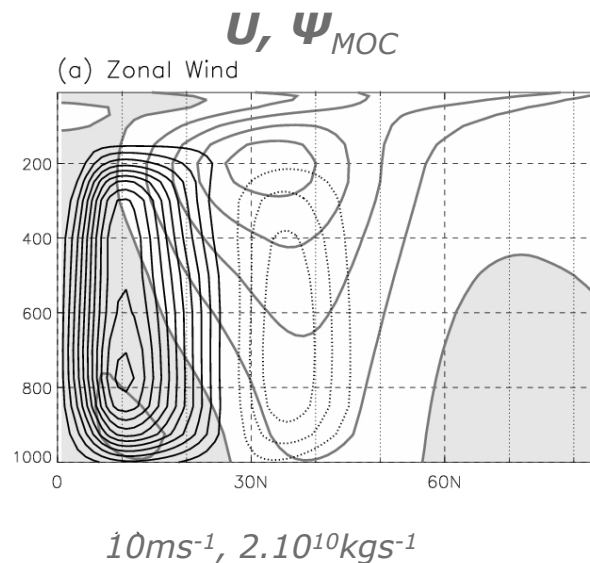


Baroclinicity  
(Eady growth rate in lower troposphere)

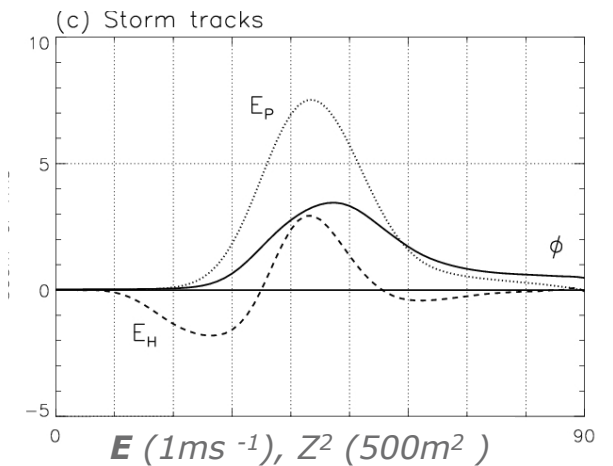


# Aqua-planet AGCM experiments

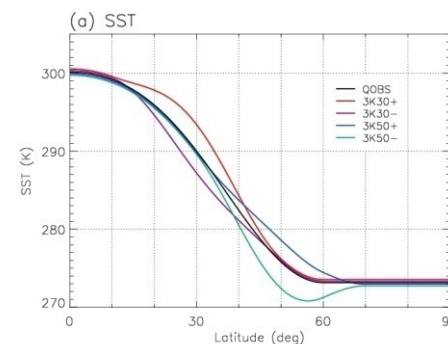
Dave Brayshaw, Mike Blackburn: Brayshaw et al (2008, 2010) JAS



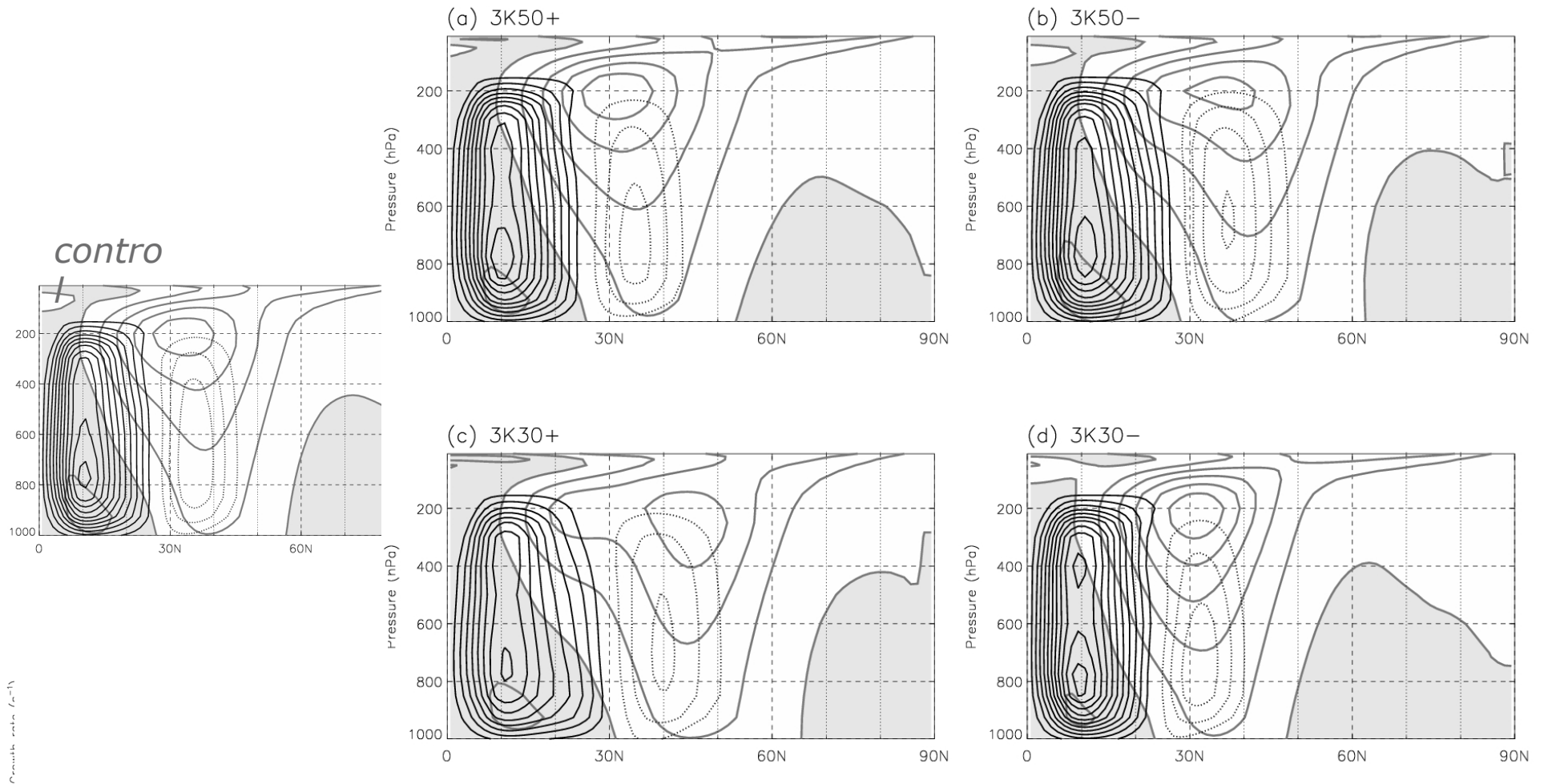
## Storm track measures



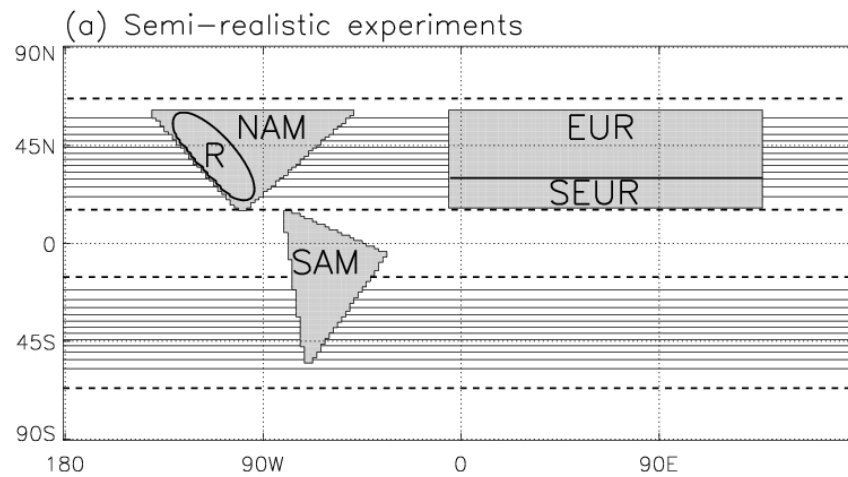
HadAM3 N96L30 ( $1.875^\circ \times 1.25^\circ$ )  
5 year runs (10 for control) after spin up  
No seasonal cycle (perpetual equinox)



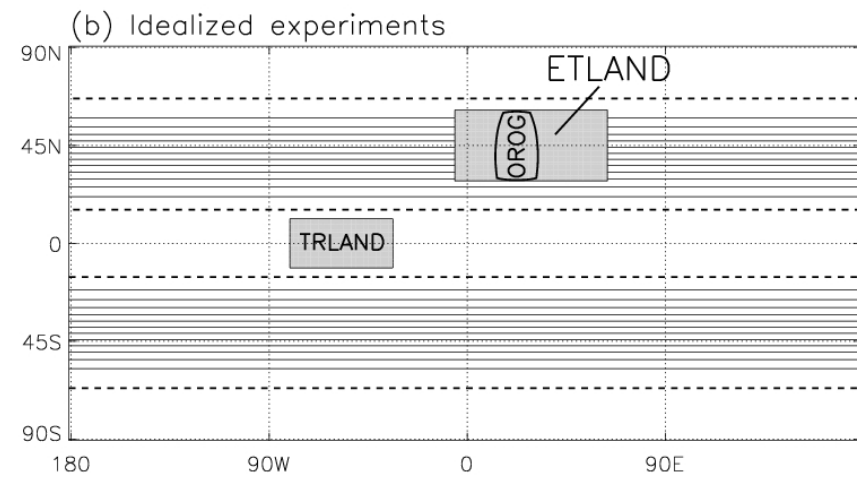
# U, $\Psi_{\text{MOC}}$ for the 4 experiments



# Continents added

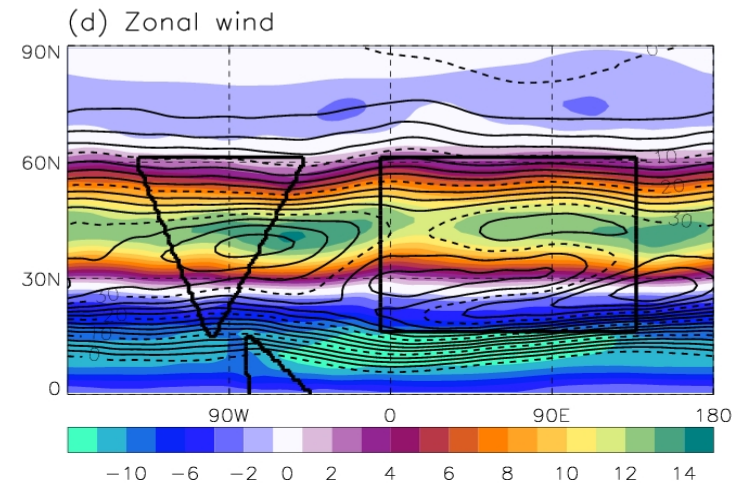
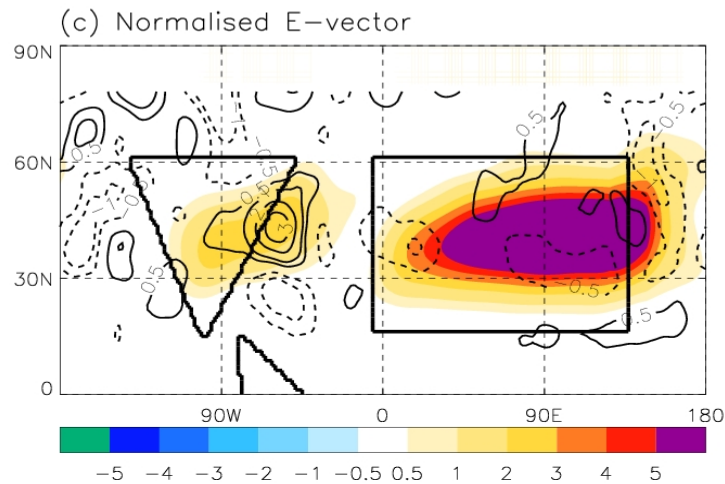
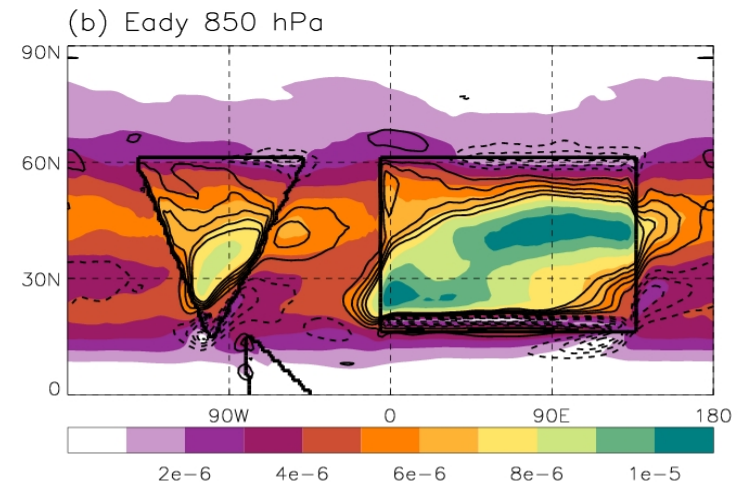
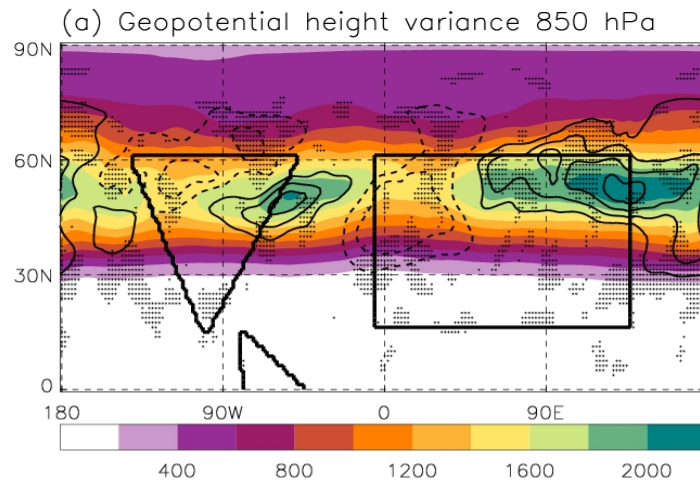


**Semi-realistic**



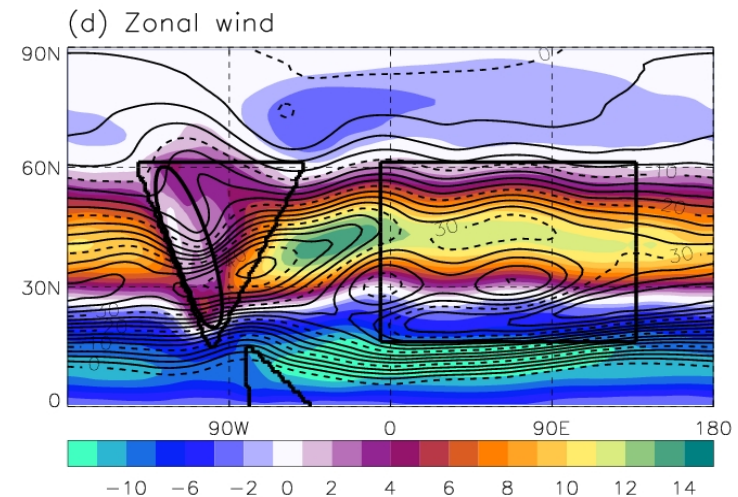
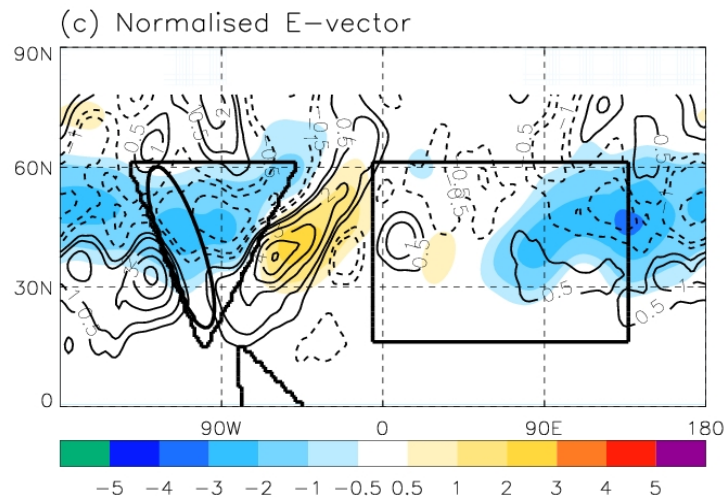
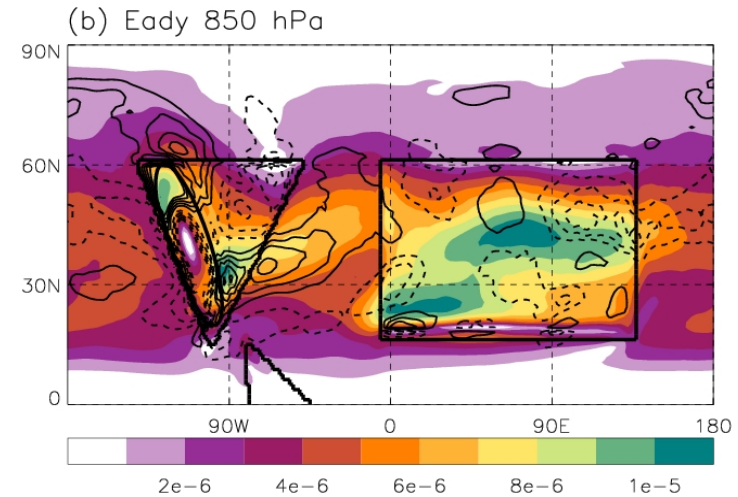
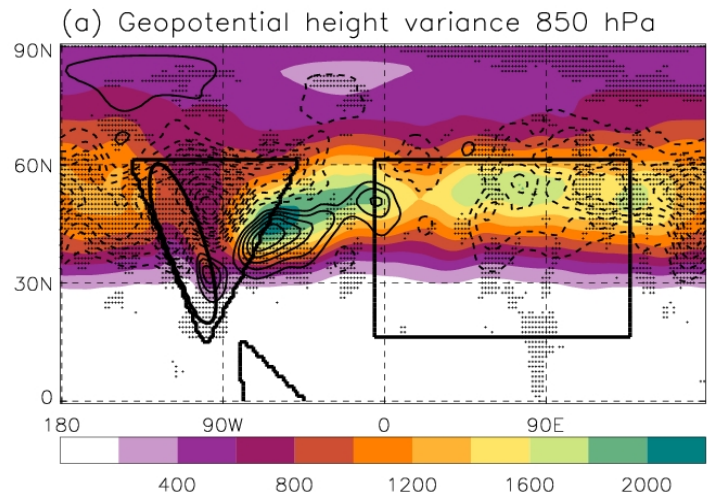
**Idealised**

# Three Continents



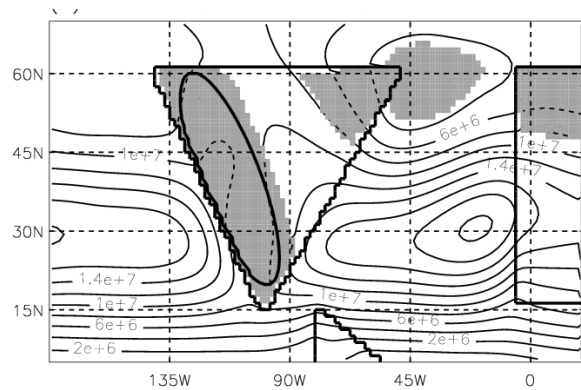


## 3 Continents + "Rockies"

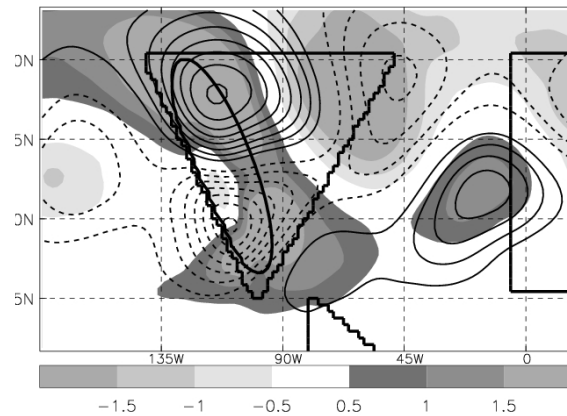




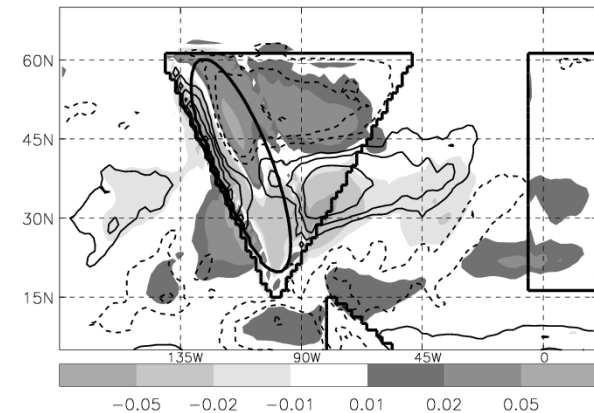
# The Impact of Adding the “Rockies”



$\psi_{1000}$



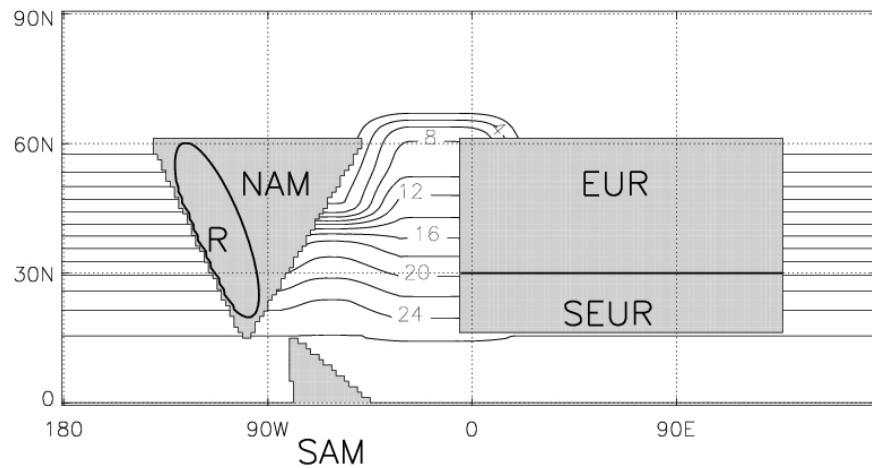
$\Delta T_{700}$ : shading  
 $\Delta \psi_{700}$ : contours



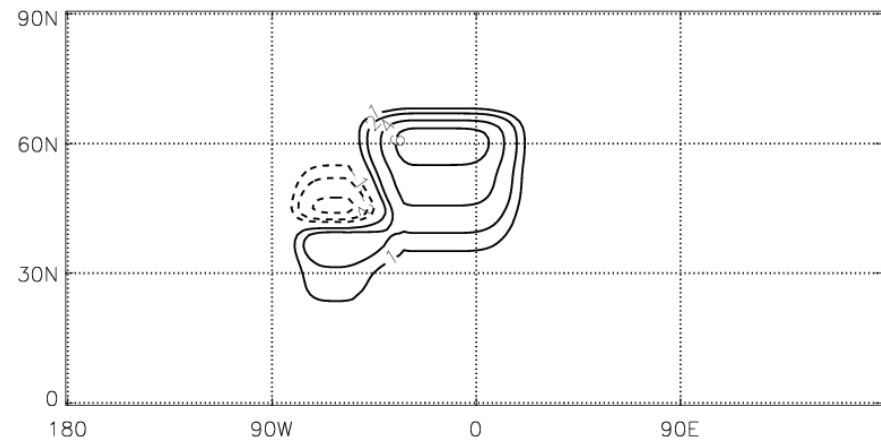
$\Delta \omega_{700}$ : shading  
 $\Delta \text{precip}$ : contours

## Adding the “Gulf Stream” & “N Atlantic Drift”

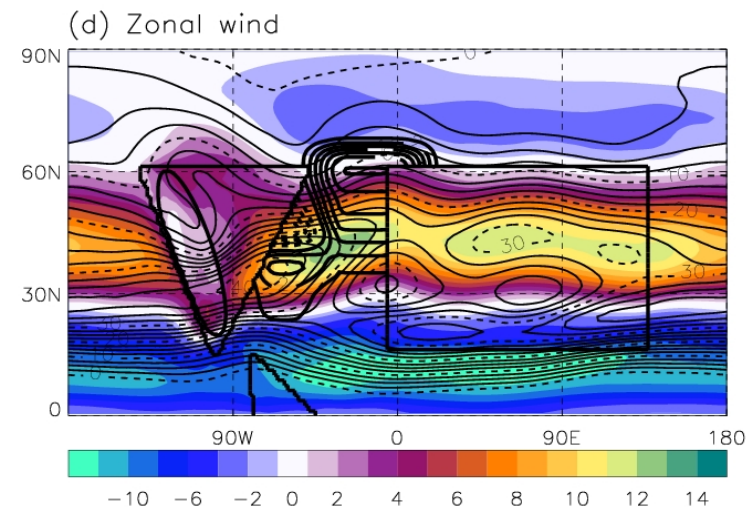
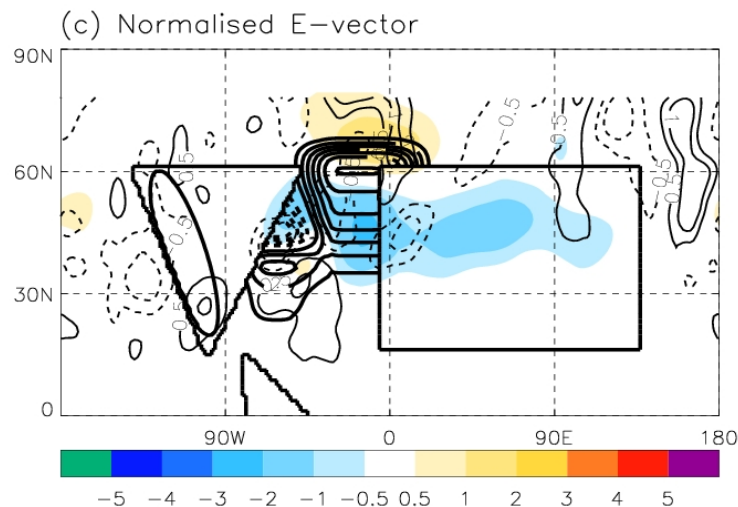
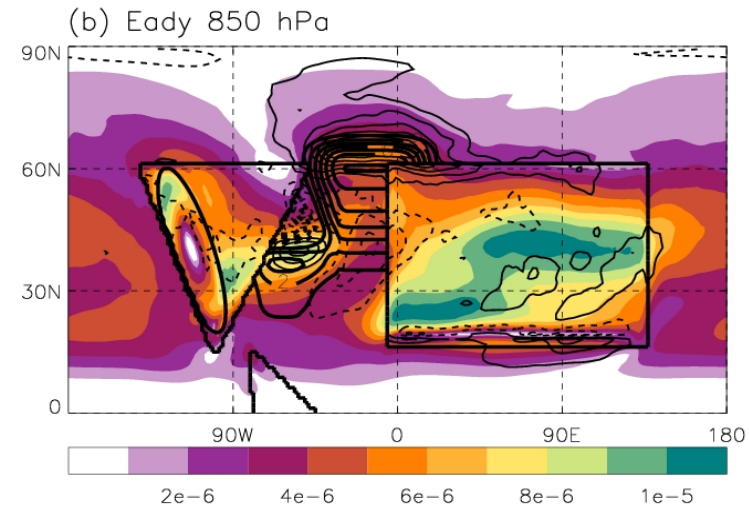
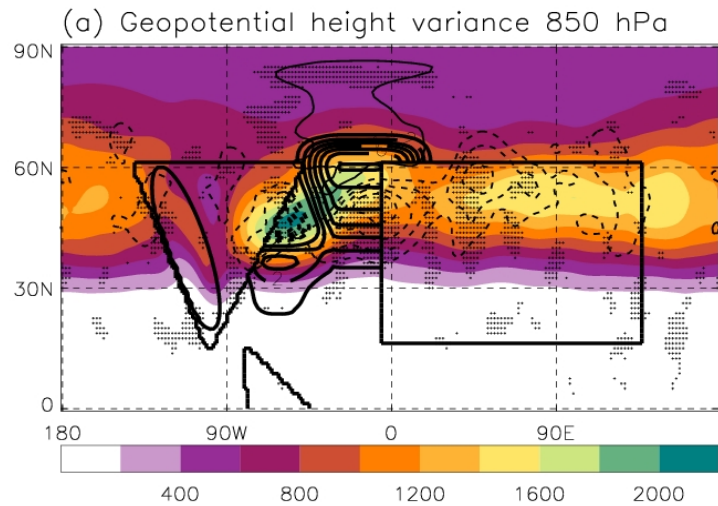
(a) With landmasses



(b) SST anomalies



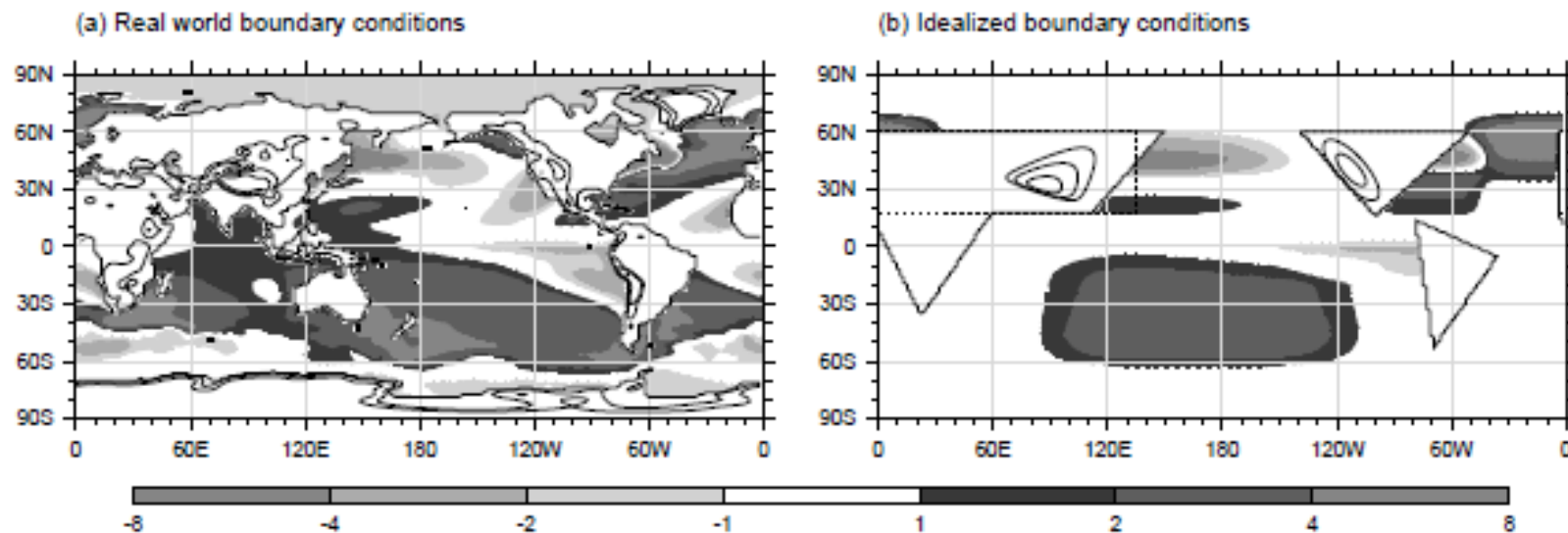
# Adding the "Gulf Stream" & the "N Atlantic Drift"



# Influence of a more realistic Eurasia & Pacific

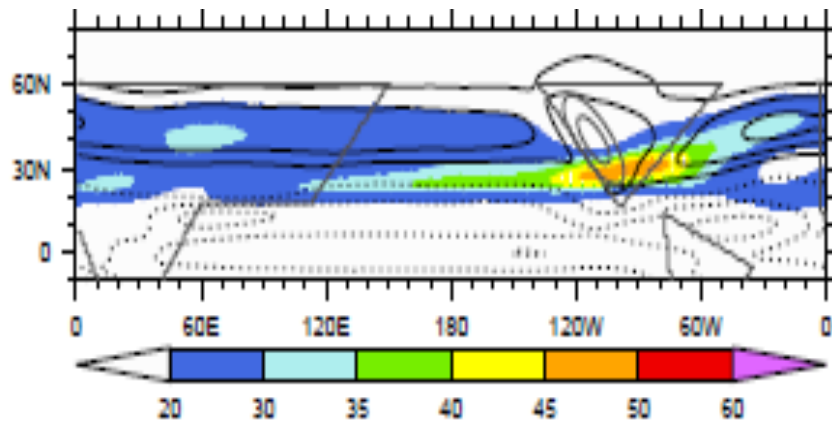
Jerome Sauliere: Sauliere et al JAS 2011

Model used: HadGAM1

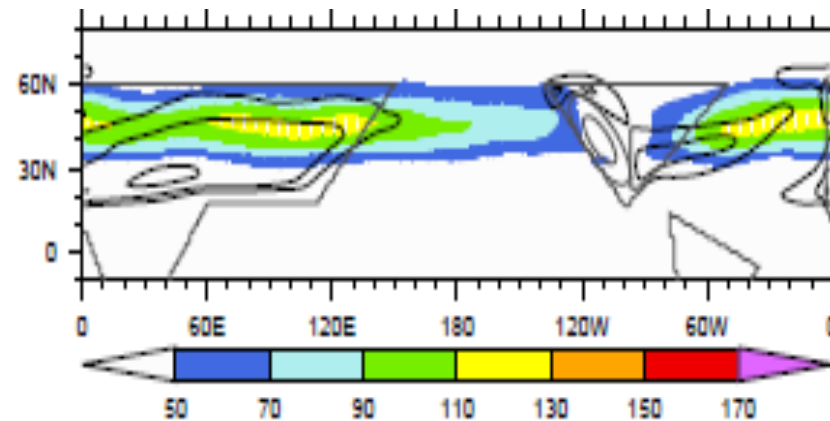


## Four Continents: better Eurasia & Africa

$U_{250}$  colours;  $U_{850}$  contours



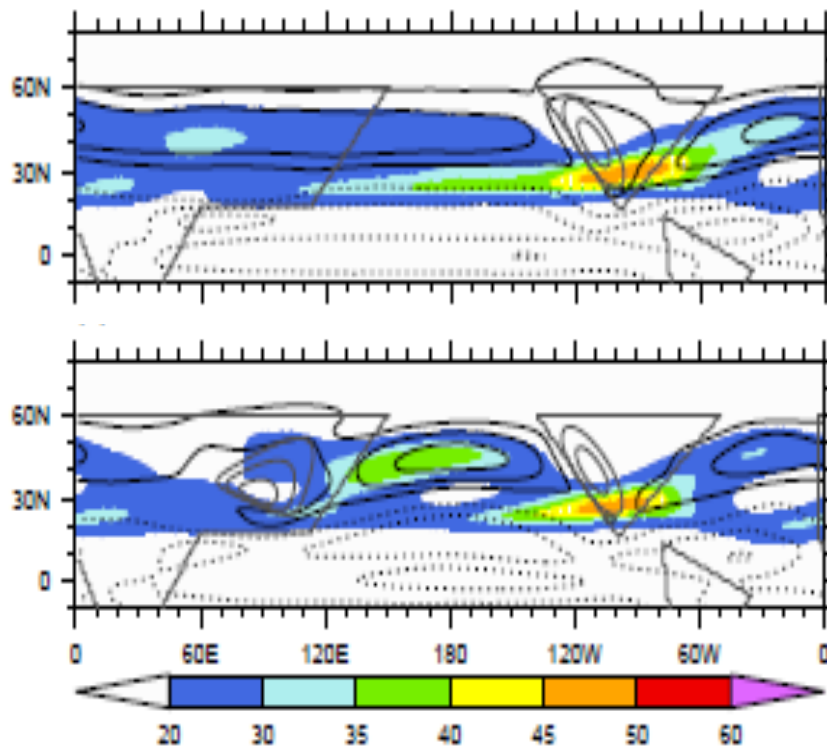
$EKE_{250}$  colours; Baroclinicity<sub>850</sub> contours



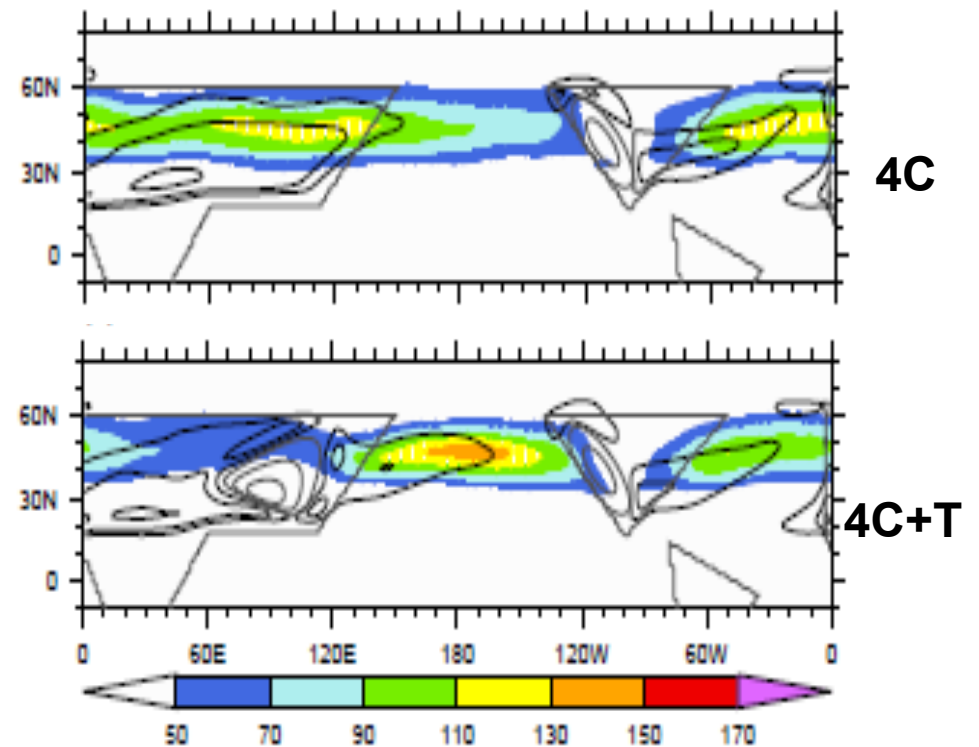


# Adding “Tibet”

$U_{250}$  colours;  $U_{850}$  contours

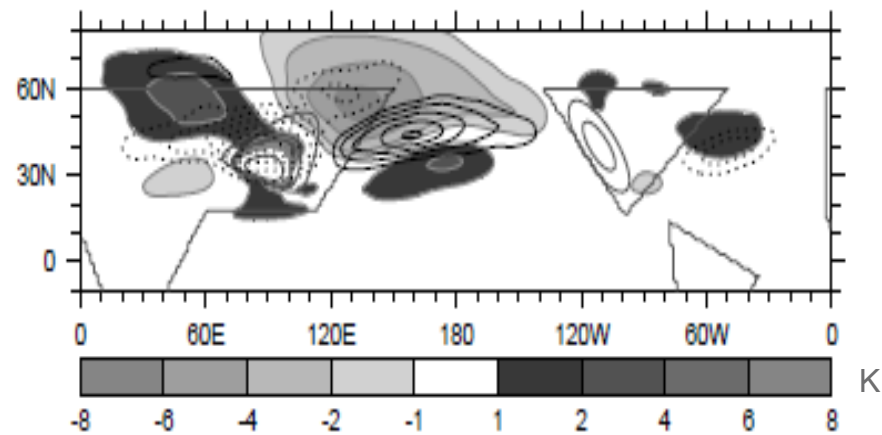


$EKE_{250}$  colours; Baroclinicity<sub>850</sub> contours



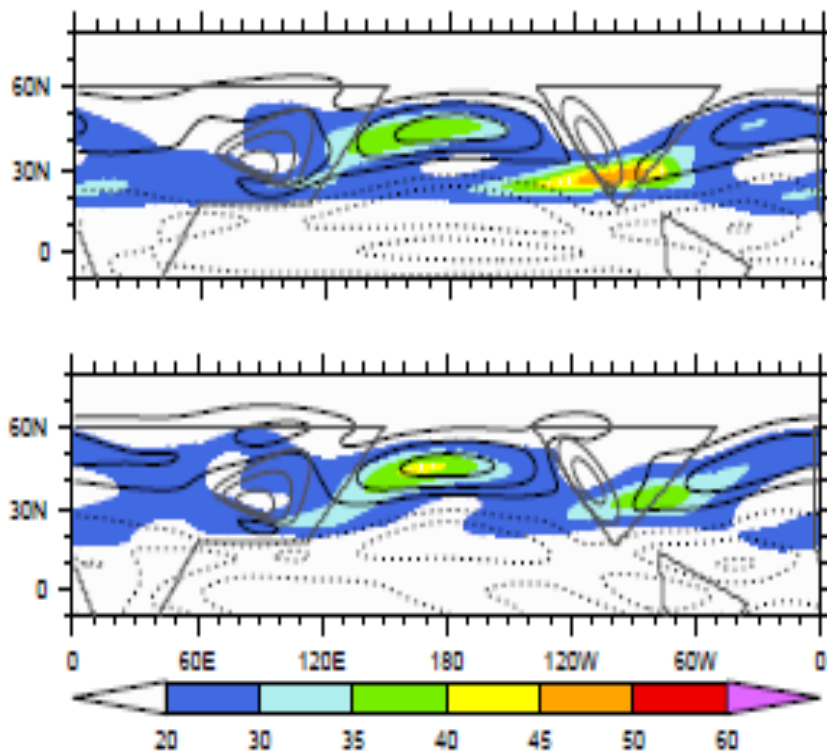
# Impact of Tibet

$\Theta_{850}$  shading;  $\omega'T'_{700}$  contours

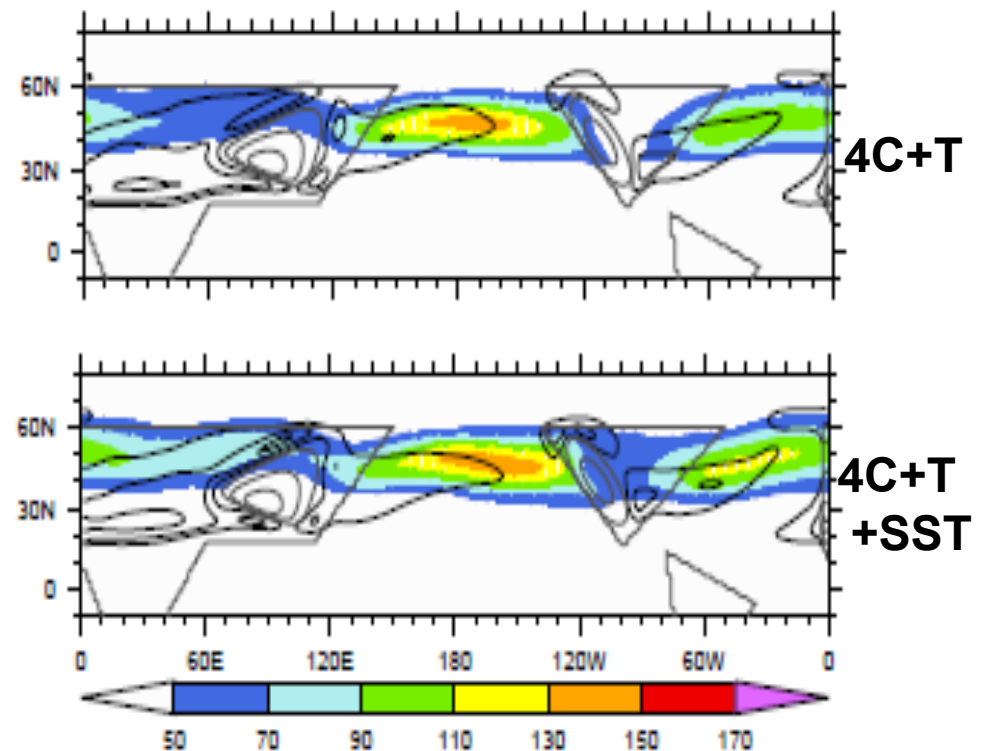


## Adding Pacific SST structure

$U_{250}$  colours;  $U_{850}$  contours



$EKE_{250}$  colours; Baroclinicity<sub>850</sub> contours



## Some aspects highlighted by the experiments

- Importance of the strength & position (wrt downward branch of Hadley Cell) of the mid-latitude SST gradient
- Small impact of extra-tropical continents: localise storm-tracks
- Strong impact of “Rockies” & “Tibet”,  
for N Atlantic & N Pacific storm-tracks, respectively
- “Gulf Stream” & “Kuroshio” produce small enhancements
- North Atlantic Drift helps stop the Atlantic storm-track continuing into “Eurasia”
- Eurasia/Pacific changes (particularly “Tibet”) influence N Atlantic storm-track

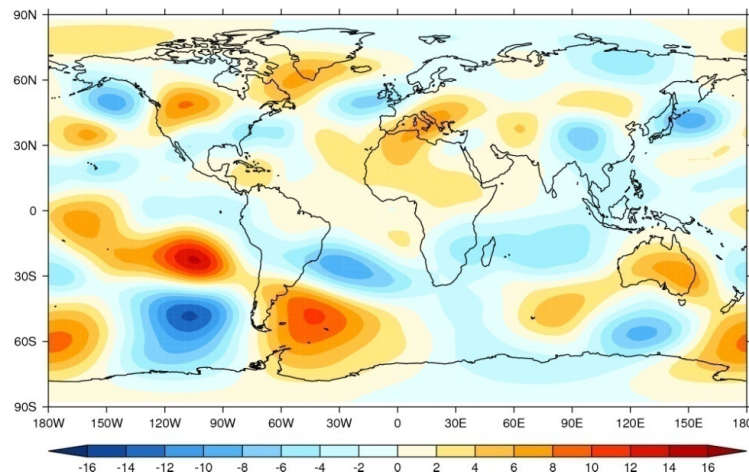
# Idealised model investigation of seasonal anomalies

Ricardo Fonseca, Mike Blackburn

Adiabatic spectral model with simple damping  
run in time-dependent mode (usually 30 days)  
Climatological mean flow maintained by forcing terms

1. **Impose tropical heating anomalies** inferred from “observation”
2. Determine extra forcing terms to maintain mean anomalous flow & investigate the role played by components of this –**Inverse Technique**
3. Relax aspects of the flow back to observed anomalies –**Relaxation**

June-July 2007

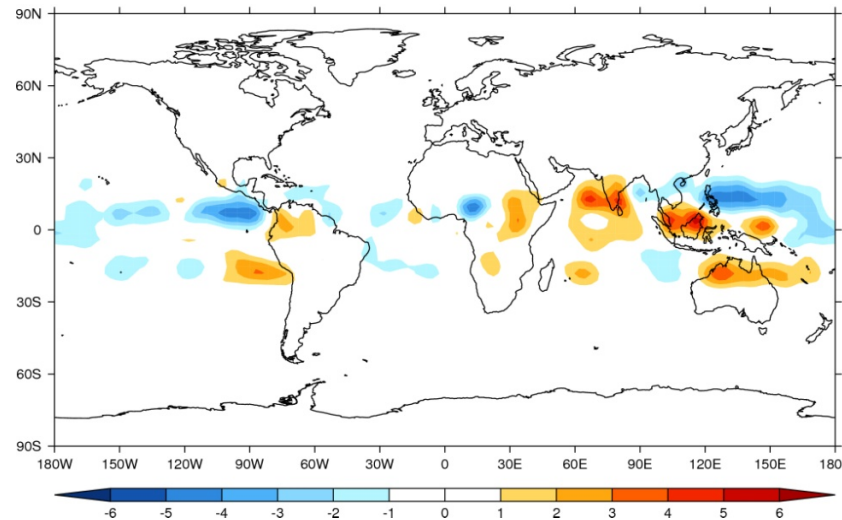


$\Psi_{200}$  anomalies

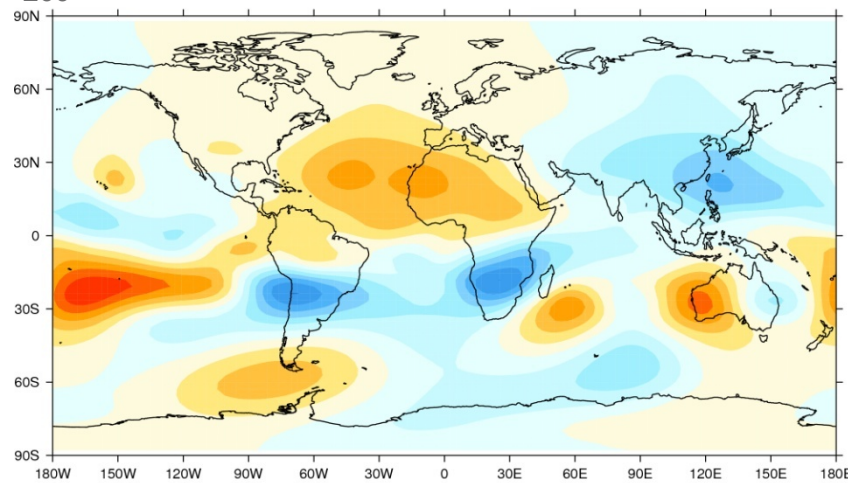


# Imposed tropical heating anomalies

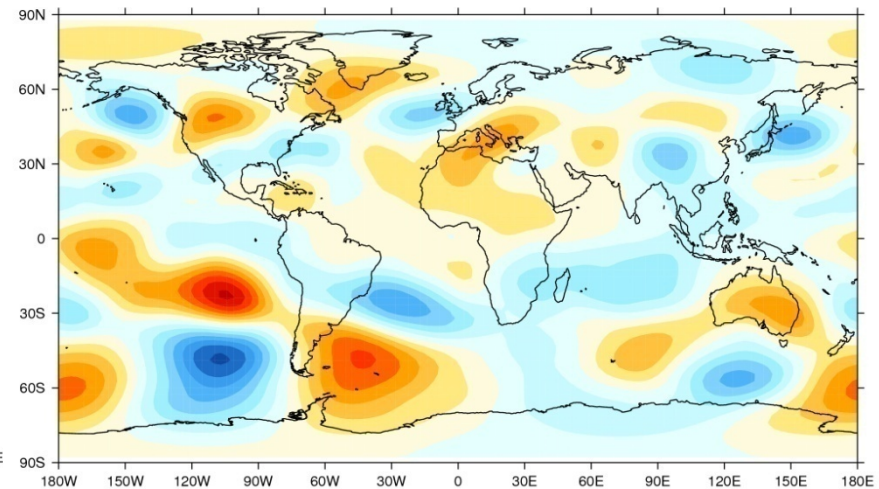
Heating at 400hPa



$\Psi_{200}$  response

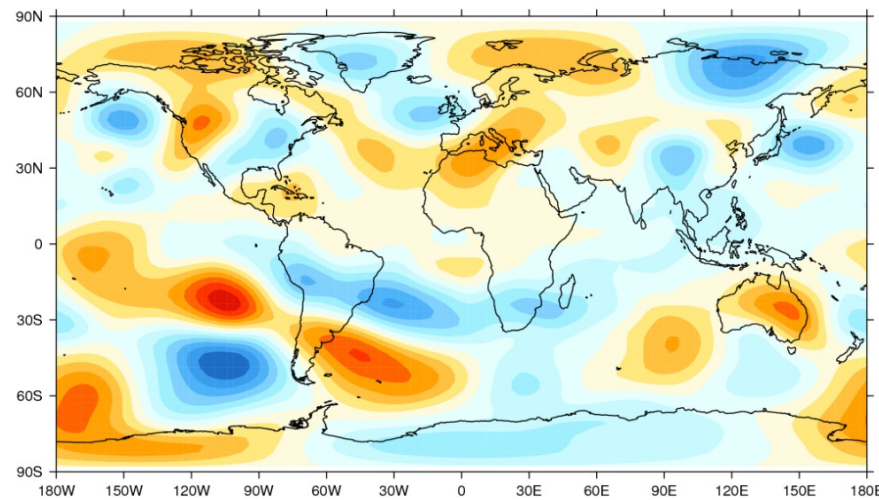


Observed

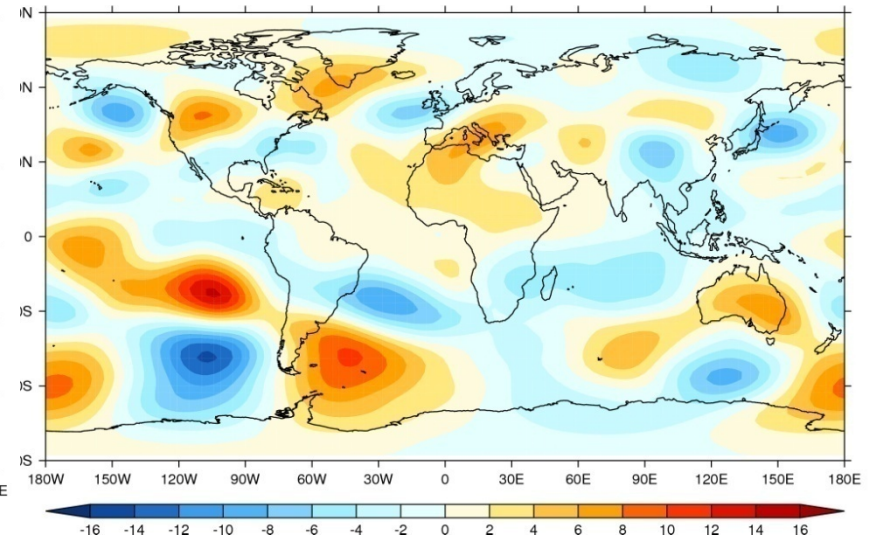


# Inverse technique: check

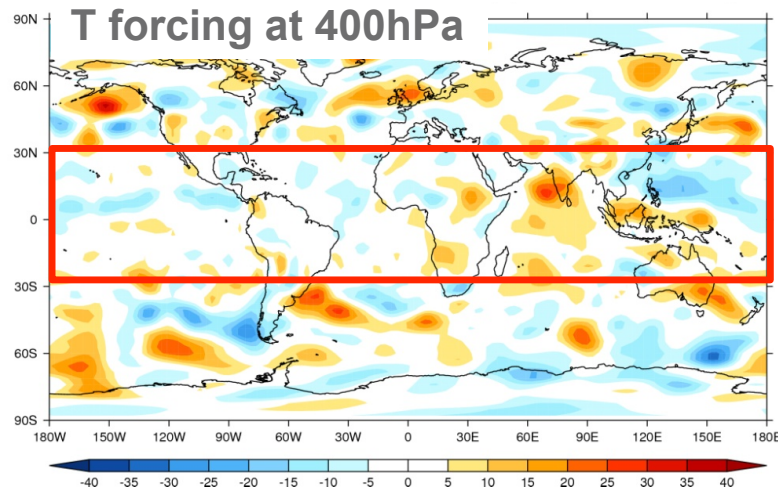
$\Psi_{200}$  response at day 30



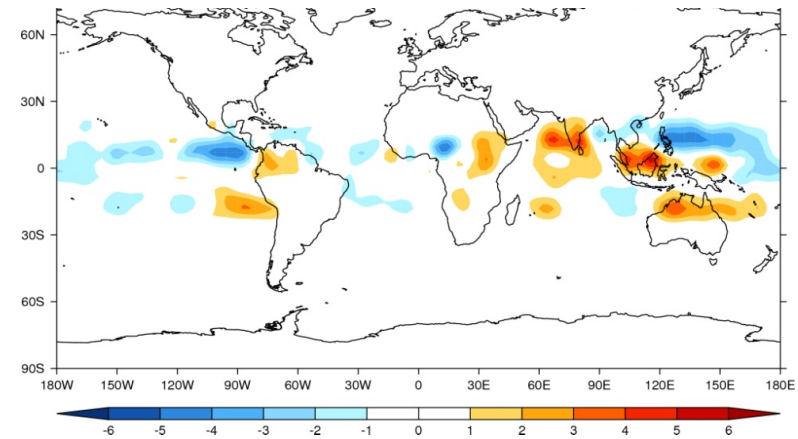
$\Psi_{200}$  observed



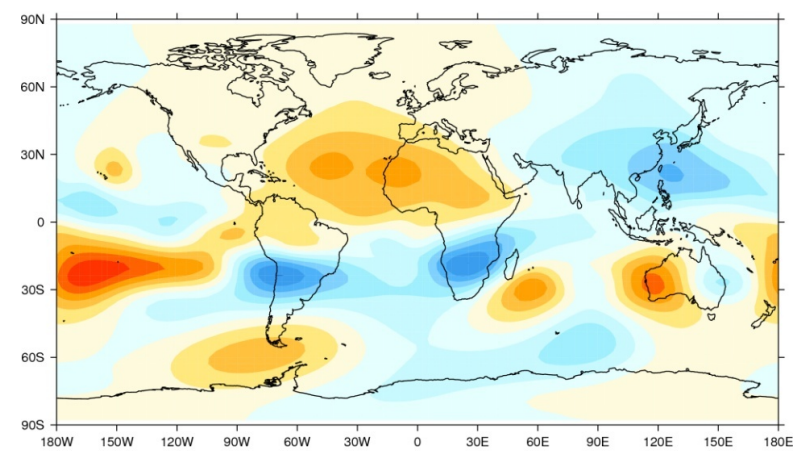
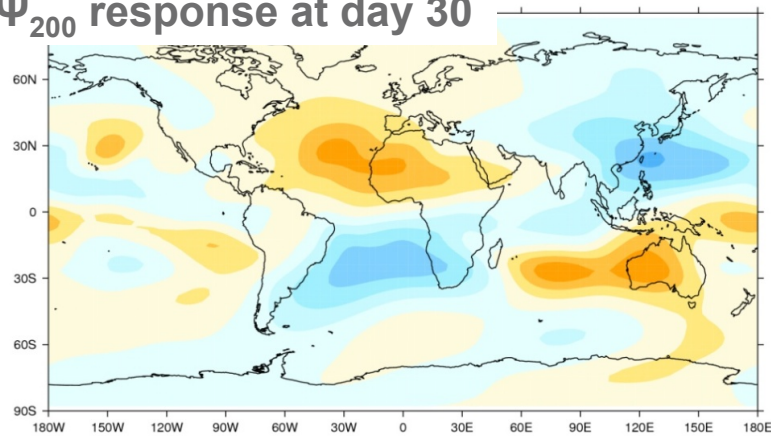
# Inverse technique: tropical T forcing



**“Observed” tropical heating anomalies**



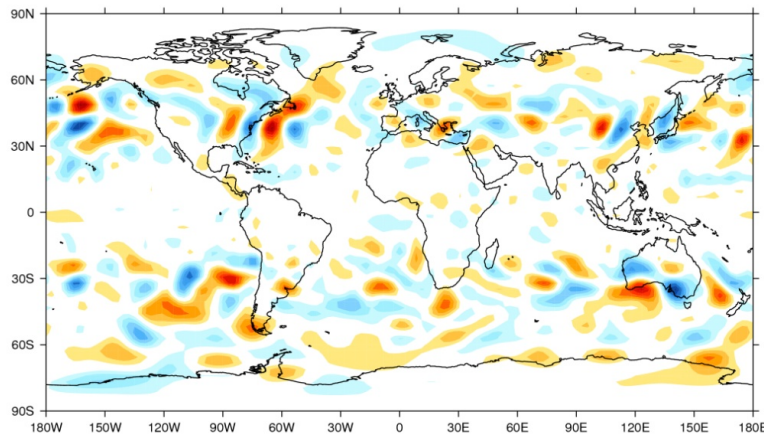
**$\Psi_{200}$  response at day 30**



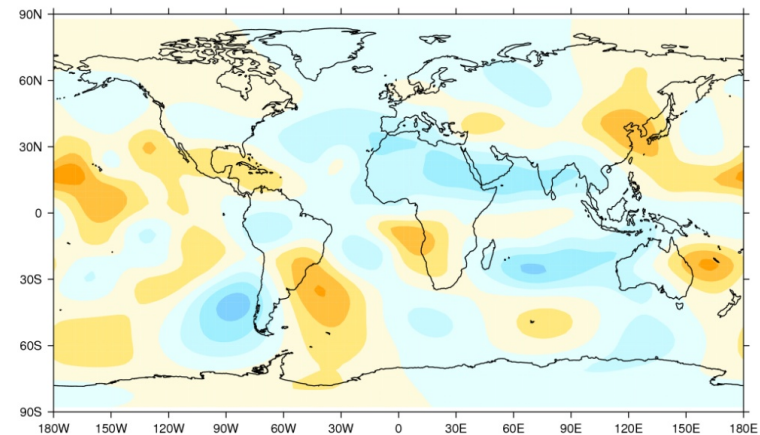


# Inverse technique: vorticity forcing

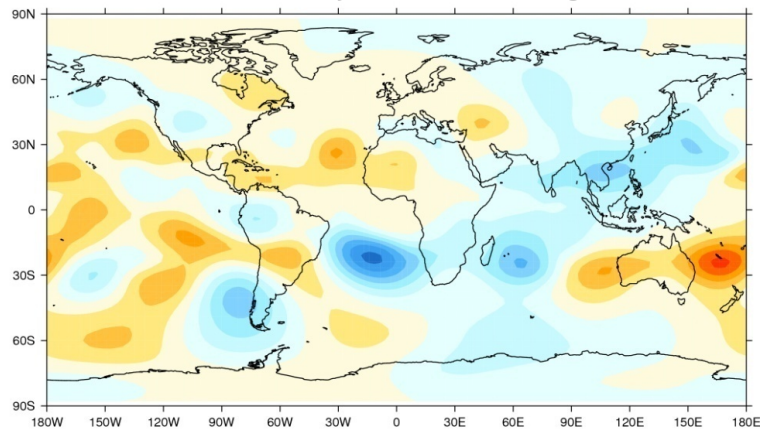
$\zeta$  Forcing at 200hPa



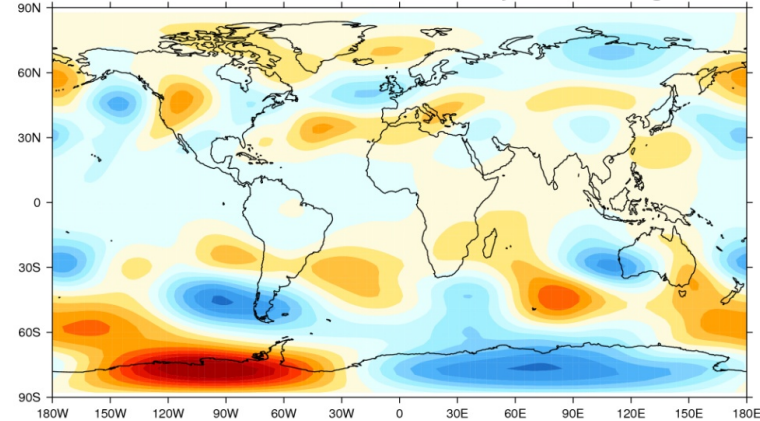
$\Psi_{200}$  response at day 30: tropical  $\zeta$  forcing



Tropical  $\zeta$  & T forcing



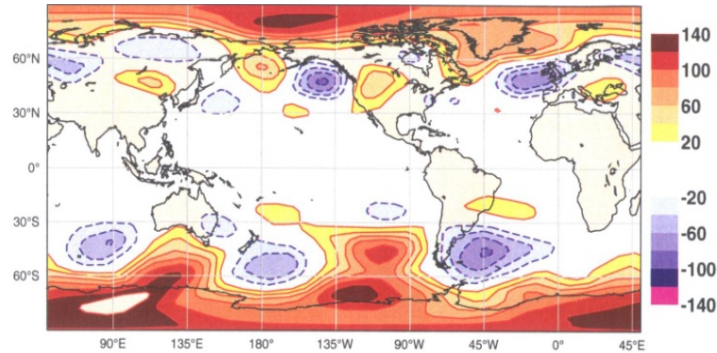
Extra-tropical  $\zeta$  forcing



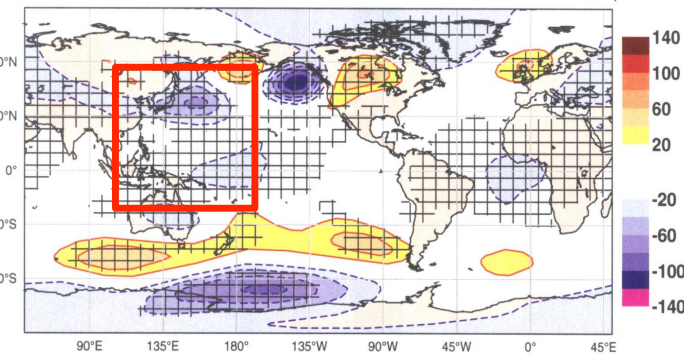
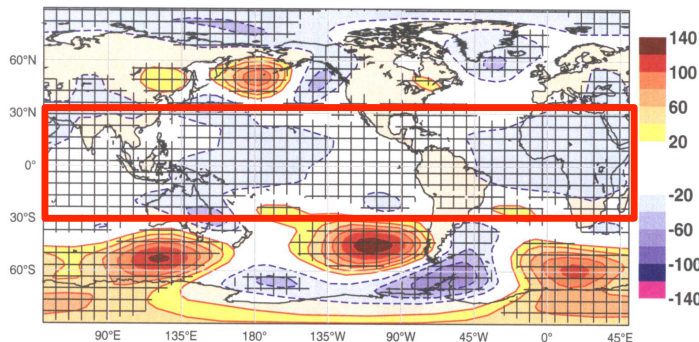
# Relaxation: comparison ECMWF ensemble

Time-scale 10h in both

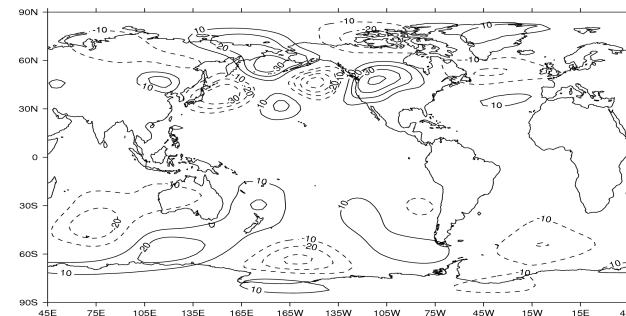
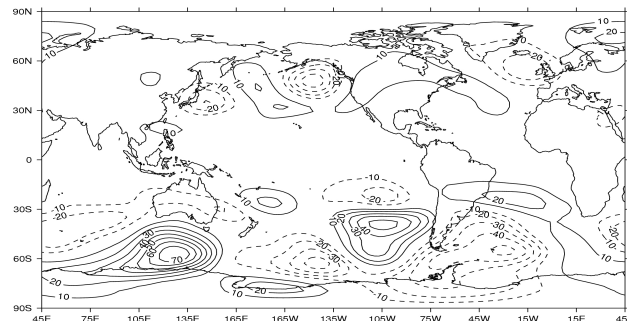
$Z_{500}$  observed



ECMWF  
MODEL  
c.i.20m  
*T Jung*

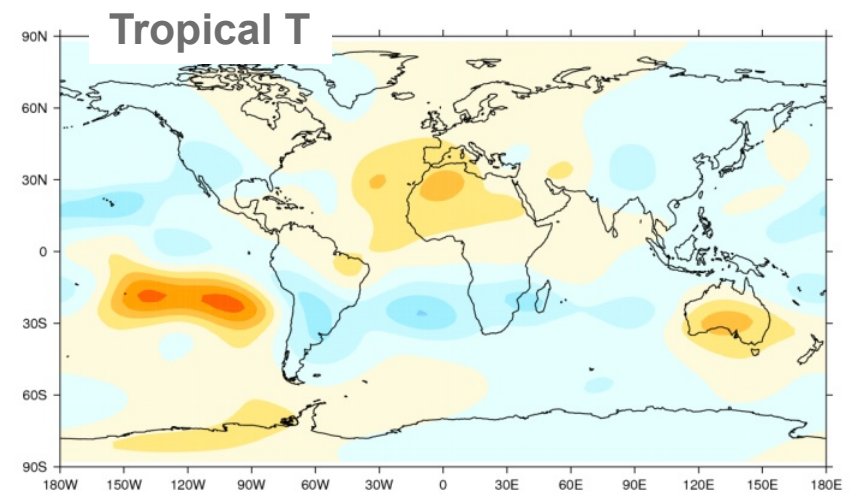
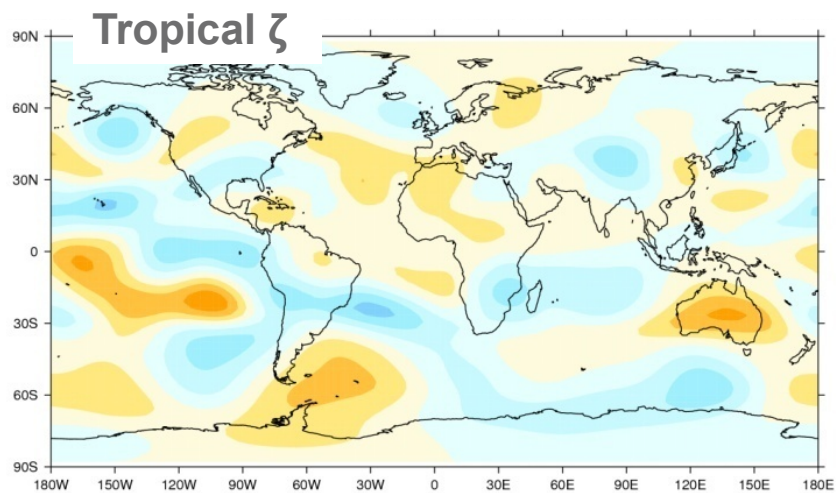
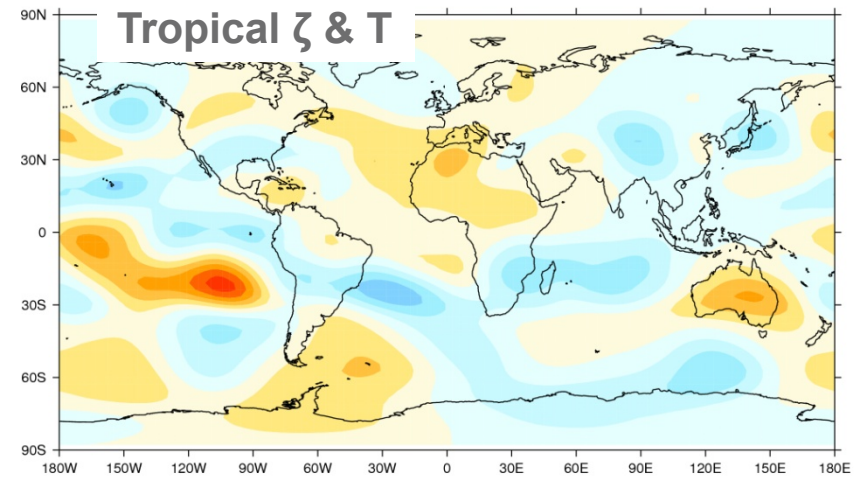
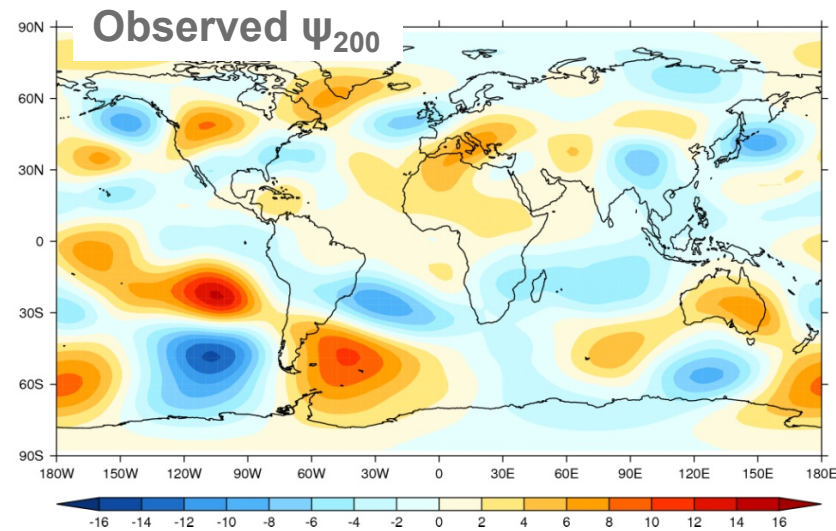


IGCM  
c.i.10m

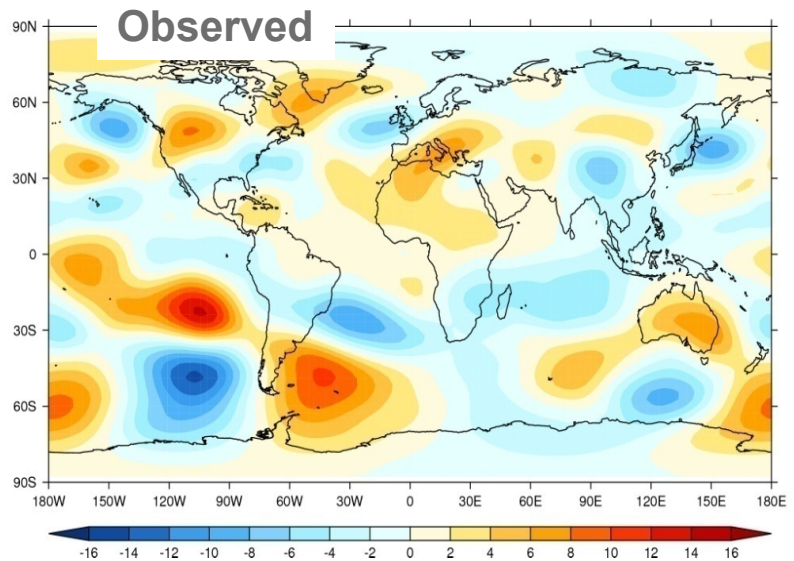
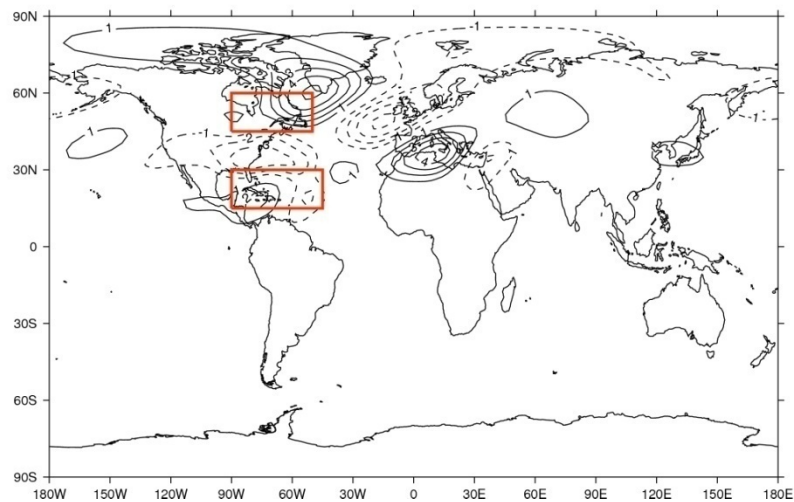
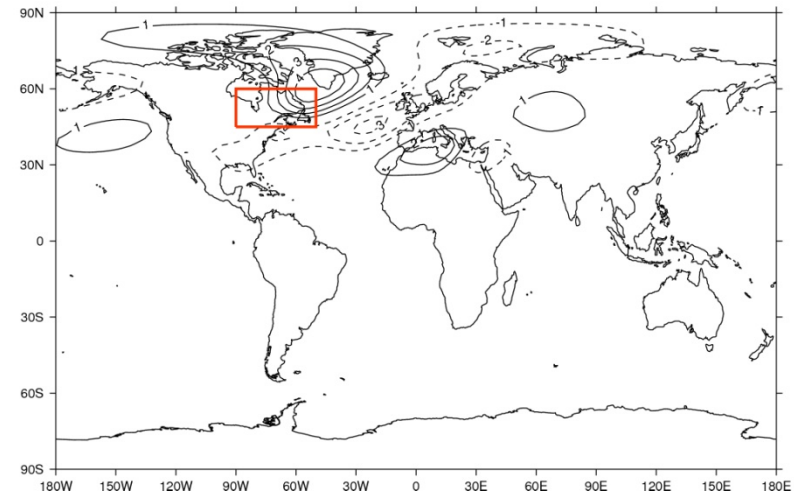
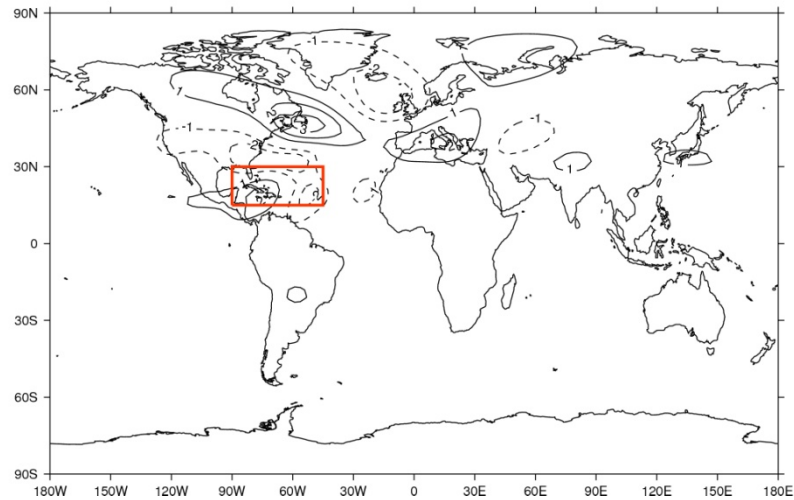




# Tropical Relaxation



## Relaxation: localised



# Conclusion

- 3 techniques with a simple time-dependent baroclinic model yield useful, complimentary insights
- They provide support for experiments/forecasting with complex models.
- Importance of tropical vorticity as well as thermal forcing
- Important for Europe in summer: Caribbean region  
upstream storm-track vorticity transients

**Long live the hierarchy of models!**