



2265-31

Advanced School on Understanding and Prediction of Earthquakes and other Extreme Events in Complex Systems

26 September - 8 October, 2011

Overview of Nuclear Techniques for Probing Earth's Interior and Earthquake
Prediction
Part 1 - Part 2

Wolfango Plastino

National Institute of Nuclear Physics Section of Roma Tre & Gran Sasso National Laboratory Italy



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Overview of Nuclear Techniques for Probing Earth's Interior and Earthquake Prediction

Part I - Part 2

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Outline

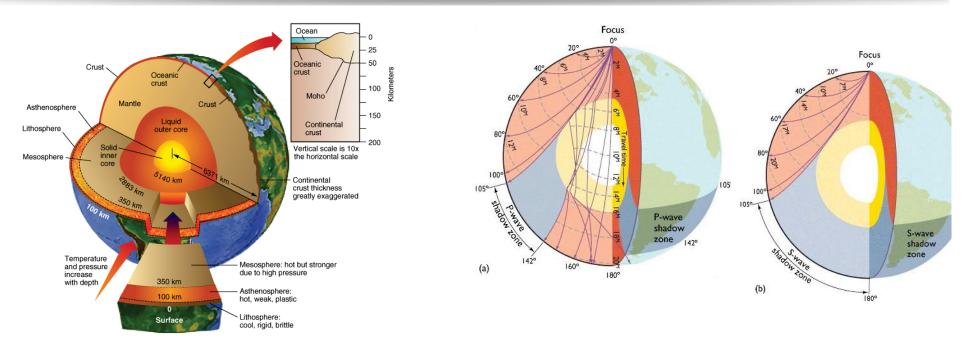
- Overview: Nuclear Physics and Radioactivity
- Earth's Interior: a view from Solid Earth Physics to Nuclear Physics
- Nuclear Physics for probing Earth's Interior and Earthquake prediction



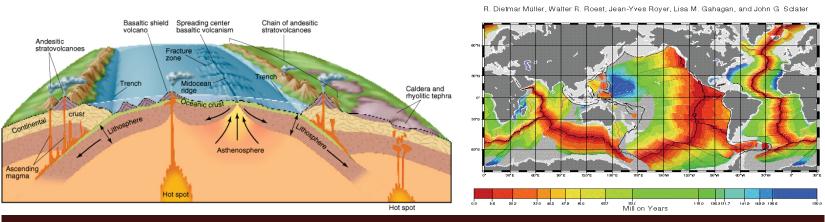




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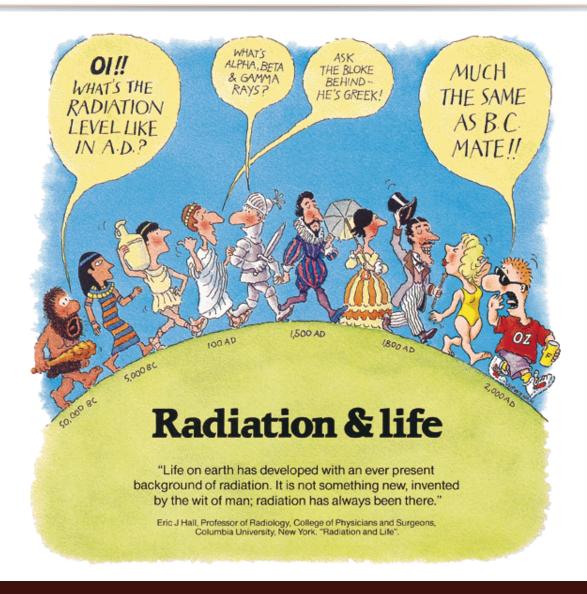
A Digital Age Map of the Ocean Floor

















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Memoria est thesaurus omnium rerum et custos

Marcus Tullius Cicero, De oratore (1, 5, 18)





Wilhelm Konrad Röntgen X Rays (Nobel Prize in Physics 1901)

Antoine Henri Becquerel – Pierre Curie – Marie Skłodowska Curie Radioactivity (Nobel Prize in Physics 1903)









Joseph John Thomson Electron (Nobel Prize in Physics 1906)

 α and β Particles (Nobel Prize in Chemistry 1908)









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Frederick Soddy Isotopes (Nobel Prize in Chemistry 1921)







Niels Bohr Structure of atoms (Nobel Prize in Physics 1922)

Werner Heisenberg Quantum Mechanics (Nobel Prize in Physics 1932)





Erwin Rudolf Josef Alexander Schrödinger Quantum Mechanics (Nobel Prize in Physics 1933)











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James Cadwick Neutron (Nobel Prize in Physics 1935)





Enrico Fermi Neutron irradiation and nuclear reactions (Nobel Prize in Physics 1938)

Ernest Orlando Lawrence Cyclotron (Nobel Prize in Physics 1939)







John Douglas Cockcroft - Ernest Thomas Sinton Walton Transmutation of atomic nuclei (Nobel Prize in Physics 1951)

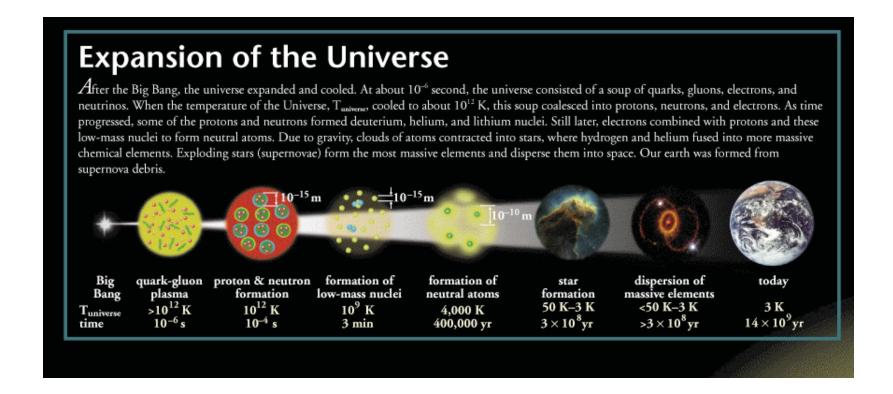




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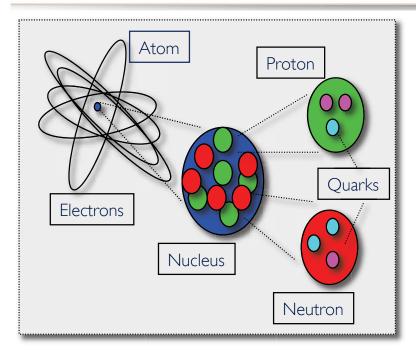


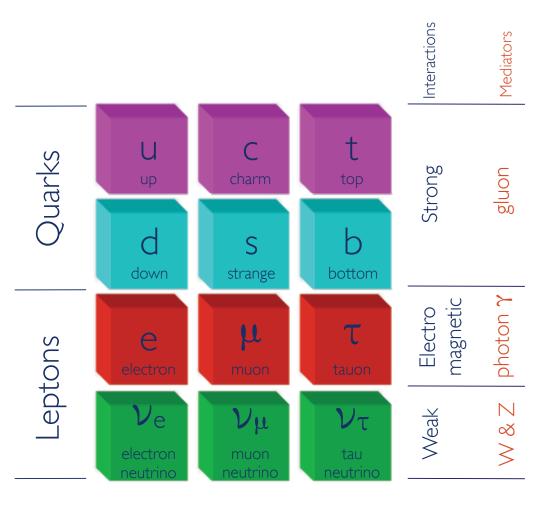


















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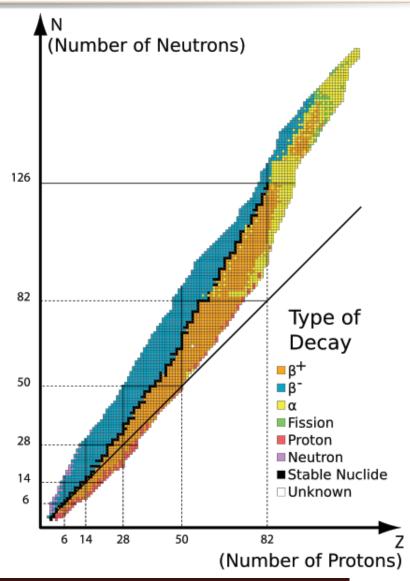


A = number of protons + number of neutrons

Z = number of protons

A - Z = number of neutrons

Number of neutrons = Mass Number – Atomic Number

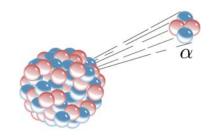


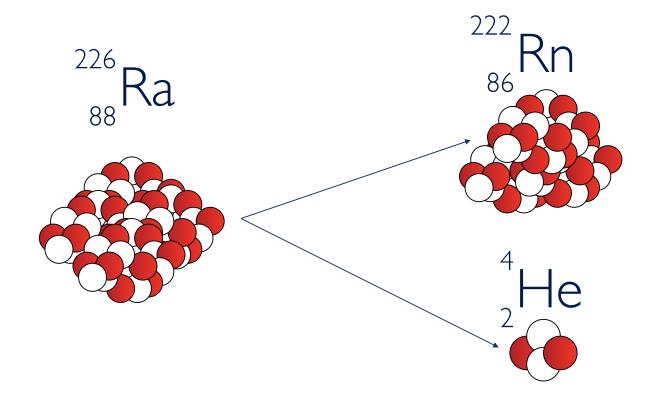




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Alpha Decay



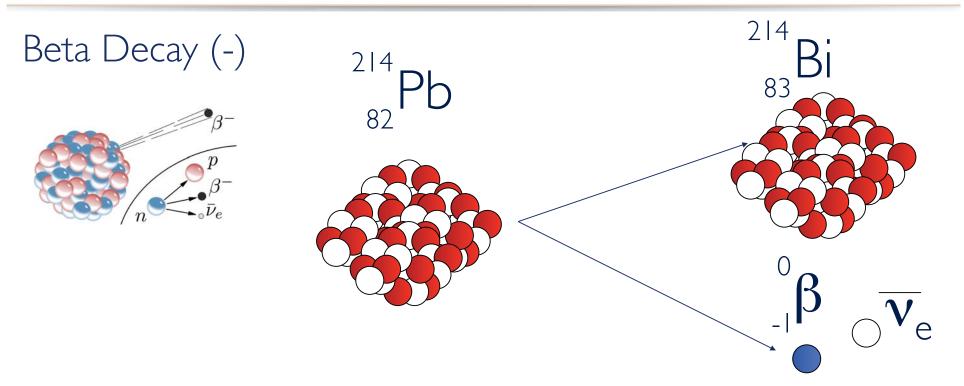


$$_{z}^{A}X \xrightarrow{A-4} Y + _{2}^{4}He$$





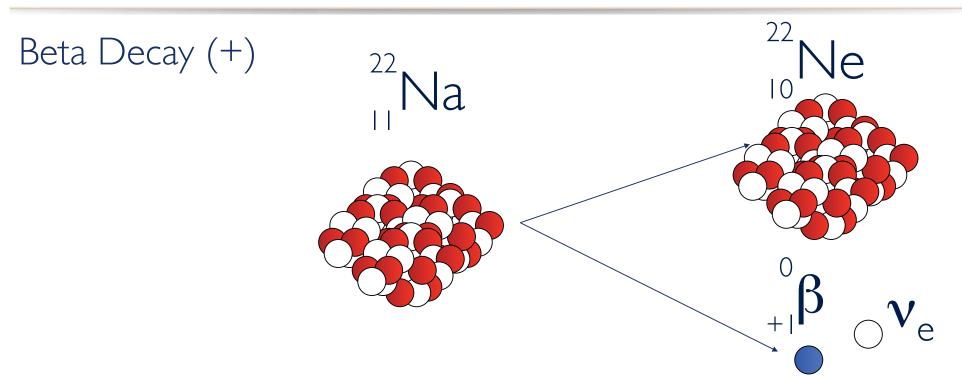




$$_{7}^{A}$$
X \xrightarrow{A} Y + $_{-1}^{0}$ β + $\overline{\mathbf{v}}_{e}$







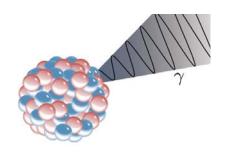
$${}_{z}^{A}X \longrightarrow {}_{z-1}^{A}Y + {}_{+1}^{0}\beta + \mathbf{v}_{e}$$

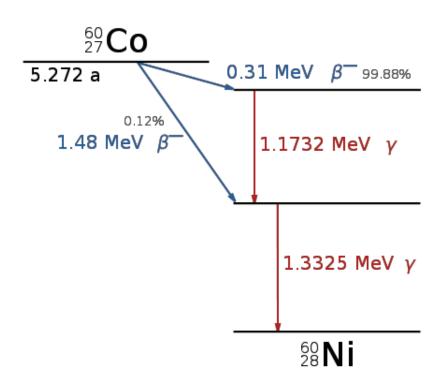




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Gamma Decay





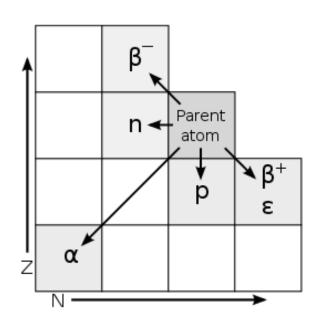
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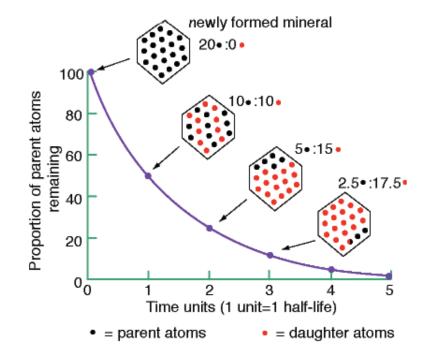




$$N(t) = N_0 e^{-t/\tau}$$

$$T_{1/2} = \tau \ln 2$$

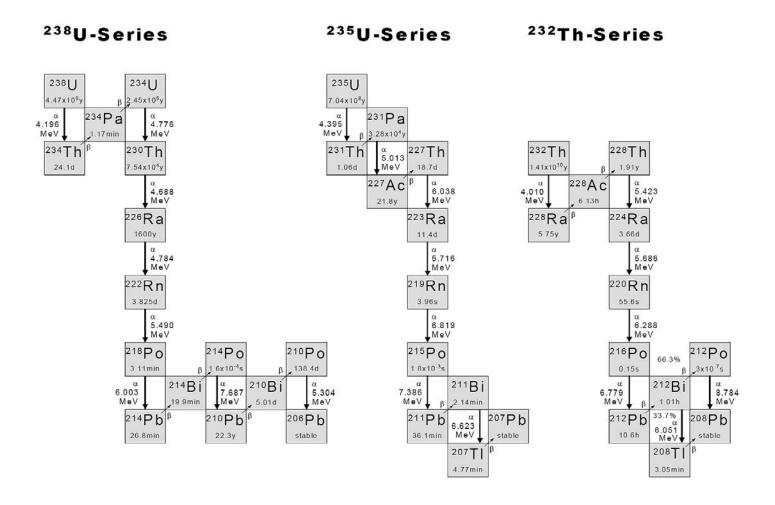








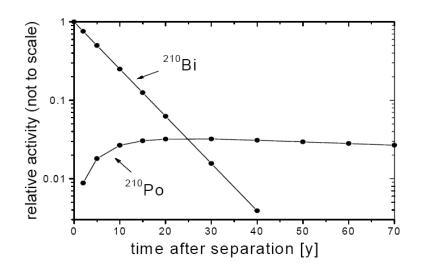
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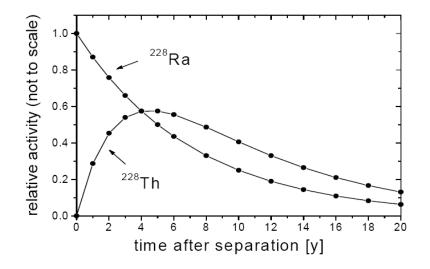


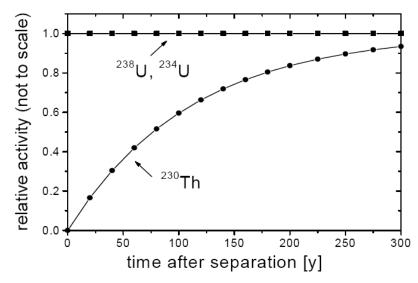


















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type of sample		$^{238}{ m U}$		²³² Th		²²⁶ Ra (1)	
	-	[ppm]	[mBq/g]	[ppm]	[mBq/g]	[ppt]	[mBq/g]
igneous	granite	2 - 10	25 - 120	5 - 30	20 - 120	0.5 - 4	25 - 120
	Gabbro	0.5 - 2	5 - 25	2 - 6	5 - 25	0.1 - 0.5	5 - 25
	Basalt	0.1 - 1	1 - 10	0.3 - 4	1 - 15	0.02 - 0.2	1 - 10
	Ultramafics	< 0.02	< 0.2	< 0.05	< 0.2	< 0.01	< 0.2
sedimentary	Shales	2 - 4	25 - 50	5 - 15	20 - 120	0.5 - 1	25 - 50
	Limestone	1 - 3	10 - 40	0 - 3	0 - 10	0.2 - 1	10 - 40
	Speleothem	1 - 3	10 - 40	0 - 3	0 - 10	0.2 - 1	10 - 40
	Coral	2 - 4	25 - 50	< 0.01	< 0.04	0.5 - 1	25 - 50
	Clay	1 - 4	10 - 50	1 - 15	5 - 60	0.2 - 1	10 - 50
water (2)	sea water (3)	3 - 4	40 - 50	< 0.01	< 0.05	0.01 - 0.1	0.5 - 5
	river water (4)	0.1 - 1	1 - 10	< 0.01	< 0.05	0.01 - 0.1	0.5 - 5







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Detector type	Principle of detection	Type of radiation	Sample source	Geometry, maximum detector efficiency	MDA [Bq/l]
α-spectrometer	semiconductive material (Si)	α, (β)	planar disc	$2\pi, 40\%(\alpha)$	4·10 ^{-4 (a)}
γ-spectrometer	semiconductive material (Ge)	γ, X-ray	direct sample measurement	2π , 10% ⁽¹⁾	5·10 ^{-2 (b)}
NaI-Detektor	solid scintillator (NaI crystal)	γ, X-ray	direct sample measurement	2π , 30% ⁽²⁾	5·10 ^{-2 (b)}
α/β proportional counter	counting gas ionization	α, β	planar disc	2π , $45(\beta)$, $30(\alpha)$	1·10 ^{-3 (a)}
grid ionization chamber	counting gas ionization	α , (β)	planar disc	$2\pi, 30\%(\alpha)$	4·10 ^{-3 (c)}
α/β liquid scint. spectrometer	liquid scintillator	α, β	scintillation cocktail	2π, 100%(α+β)	1·10 ^{-2 (d)}

⁽¹⁾ valid for a 30% efficiency intrinsic Ge detector (relative to 3x3' NaI) using a 3 cm planar source measured on the detector surface and for a photon energy of 200 keV





valid for a 3x3' Ø NaI-Detector using a 3 cm planar source measured on the detector surface and for a photon energy of 200 keV

⁽a) radiochemical separation (sorption onto MnO₂-coated discs) from a 1 liter sample

⁽b) measurement of the ²²²Rn progenies ²¹⁴Pb, ²¹⁴Bi in gas tight containers (6 liter simultaneously)

evaporation of 100 ml water in a large surface (20 cm \varnothing) sample holder

⁽d) extraction of ²²²Rn from a 100 ml water sample into an organic cocktail using a separation funnel



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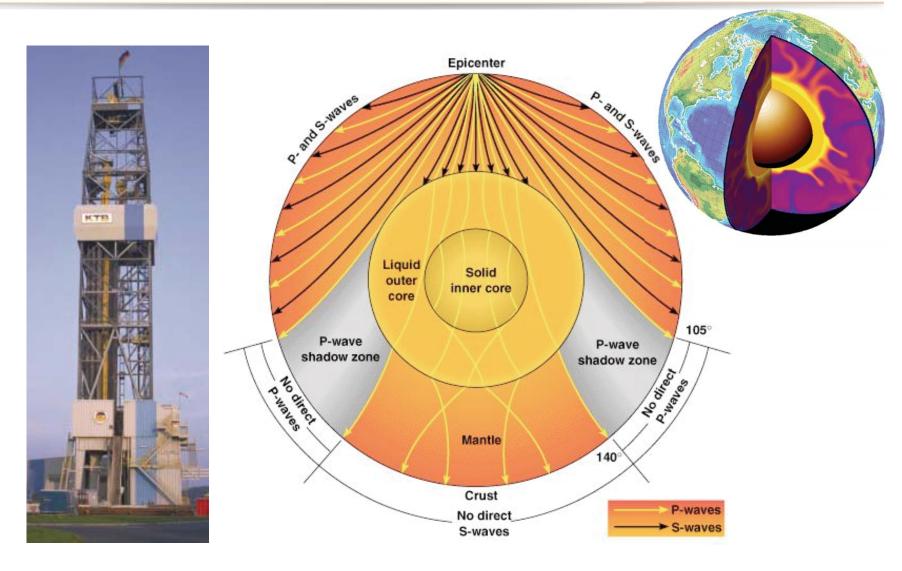
isotope	origin	mode of decay	mode of detection	detection limit [Bq/g]	detection limit [g/g]
²³⁸ U	²³⁸ U-series	α, (γ)	α-spectrometry	5.10-5	4·10 ⁻⁹
$^{235}{ m U}$	²³⁵ U-series	α, γ	α -spectrometry	5.10-5	6·10 ⁻¹⁰
²³⁴ U	²³⁸ U-series	α, (γ)	α -spectrometry	5·10 ⁻⁵	2·10 ⁻¹³
^{236}U	spike isotope	α, (γ)	α -spectrometry	5·10 ⁻⁵	8·10 ⁻¹¹
^{233}U	spike isotope	α , (γ)	α -spectrometry	5·10 ⁻⁵	1·10 ⁻¹³
$^{232}{ m U}$	spike isotope	α , (γ)	α -spectrometry	5.10-5	6·10 ⁻¹⁷
²³¹ Pa	²³⁵ U-series	α, (γ)	α-spectrometry	5·10 ⁻⁵	3·10 ⁻¹⁴
²³³ Pa	spike isotope	β, γ	γ-spectrometry	$1 \cdot 10^{-1}$	$1 \cdot 10^{-16}$
²³² Th	²³² Th-series	α, (γ)	α-spectrometry	5·10 ⁻⁵	1·10 ⁻⁸
²³⁰ Th	²³⁸ U-series	α , (γ)	α -spectrometry	5·10 ⁻⁵	7·10 ⁻¹⁴
²²⁸ Th	²³² Th-series	α , (γ)	α -spectrometry	5.10-5	2.10-18
²²⁹ Th	spike isotope	α , (γ)	α -spectrometry	5.10-5	5·10 ⁻¹⁴
²²⁸ Ra	²³² Th-series	β	proportional counting	1·10 ⁻³	1·10 ⁻¹⁶
²²⁶ Ra	²³⁸ U-series	α (γ)	α -spectrometry	5·10 ⁻⁵	1·10 ⁻¹⁵
²²⁶ Ra	²³⁸ U-series	α , (γ)	α/β - LSC	5·10 ⁻⁴	$1 \cdot 10^{-14}$
²²⁴ Ra	²³² Th-series	α , (γ)	α -spectrometry	$1 \cdot 10^{-4}$	2.10-20
²²³ Ra	²³⁵ U-series	α , (γ)	α -spectrometry	5·10 ⁻⁵	3·10 ⁻²⁰
²¹⁰ Pb	²³⁸ U-series	β, (γ)	proportional counting	1·10 ⁻³	3·10 ⁻¹⁶
²¹⁰ Pb	²³⁸ U-series	β , (γ)	α/β -LSC	5·10 ⁻³	2·10 ⁻¹⁵
²¹⁰ Po	²³⁸ U-series	α	α-spectrometry	5·10 ⁻⁵	3·10 ⁻¹⁹
²⁰⁹ Po	spike isotope	α	α -spectrometry	5·10 ⁻⁵	6·10 ⁻¹⁷
²⁰⁸ Po	spike isotope	α	α -spectrometry	5·10 ⁻⁵	3·10 ⁻¹⁸

...to be continued





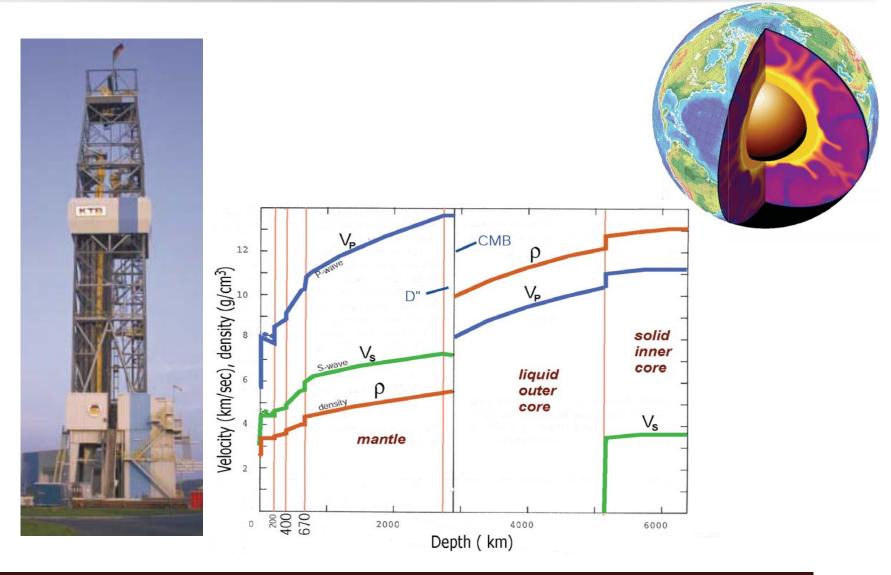
















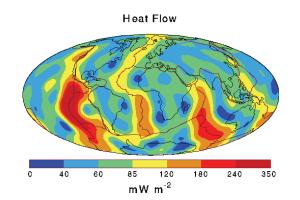


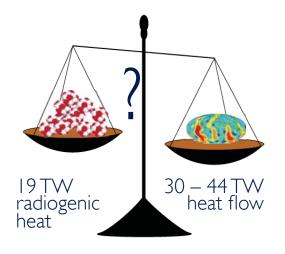
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- "Energetics of the Earth and the missing heat source mistery" *
- Heat flow from the Earth is the equivalent of some 10000 nuclear power plants

$$H_{Earth} = (30 - 44) TW$$

- The BSE canonical model, based on cosmochemical arguments, predicts a radiogenic heat production ~ 19 TW:
- ~ 9 TW estimated from radioactivity in the (continental) crust
- ~ IOTW supposed from radioactivity in the mantle
- ~ 0 TW assumed from the core
- Unorthodox or even heretical models have been advanced...





* D. L. Anderson (2005), Technical Report, www.MantlePlume.org





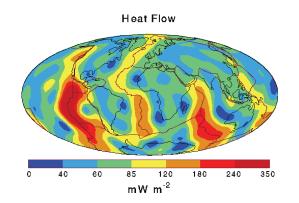


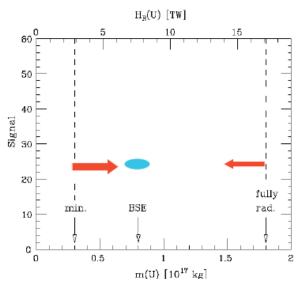
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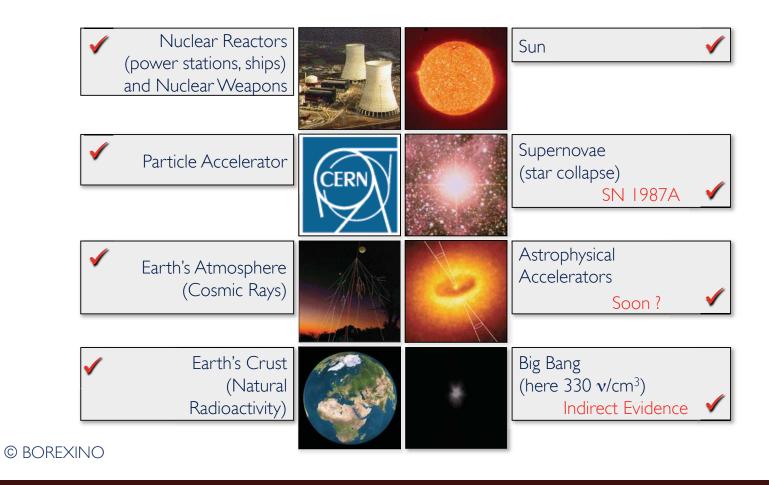
				Interactions	Mediators
Quarks	U up	C charm	top	Strong	gluon
O'	d	S strange	bottom		- 00
-eptons	electron e	μ muon	T tauon	Electro	photon γ
Lept	Ve electron neutrino	ν _μ muon neutrino	ν τ tau neutrino	Weak	× × ×

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Where do Neutrinos come from?





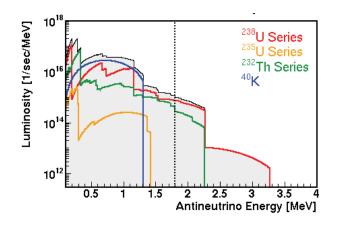


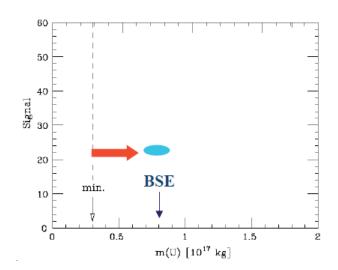


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Decay	$T_{1/2}$
	$[10^9 \text{ yr}]$
$^{238}{ m U} ightarrow ^{206}{ m Pb} + 8\ ^4{ m He} + 6e + 6ar{ u}$	4.47
$^{232}{ m Th} ightarrow ^{208}{ m Pb} + 6 ^{4}{ m He} + 4e + 4\bar{\nu}$	14.0
$^{40}\text{K} \to ^{40}\text{Ca} + e + \bar{\nu} \ (89\%)$	1.28







Earth emits (mainly) antineutrinos (106 cm⁻² s⁻¹) whereas Sun shines in neutrinos

A fraction of geo-neutrinos from U and Th (not from 40 K) are above threshold for inverse β on protons







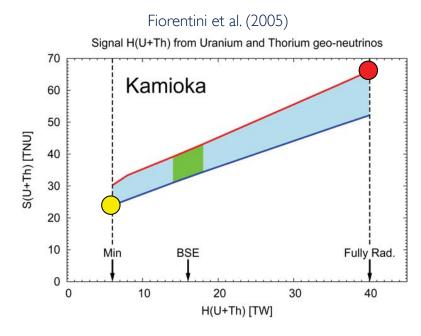
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Geo-neutrino signal and radiogenic heat from the Earth

region allowed by BSE: signal between 31 and 43 TNU

region containing all models consistent with geochemical and geophysical data

- O U and Th measured in the crust implies a signal at least of 24 TNU
- Earth energetics implies
 the signal does not exceed 62
 TNU © BOREXINO



The graph is site dependent:

- ✓ the "slope" is universal
- ✓ the intercept depends on the site
 (crust effect)
- ✓ the width depends on the site (crust effect)

 ITNU = one event per 10³² free protons per year

 Verification

 Output

 Description

 Output

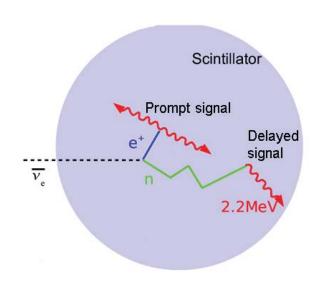
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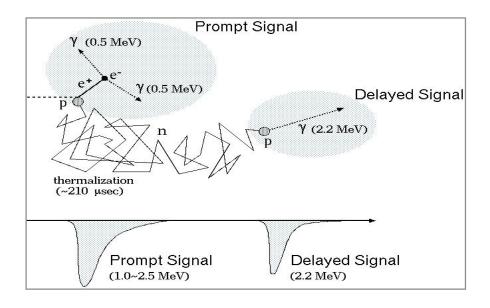






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...to be continued

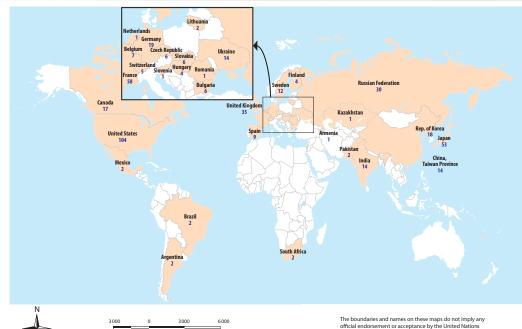
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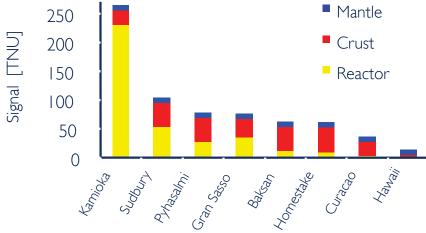




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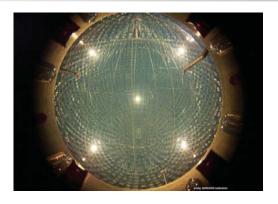








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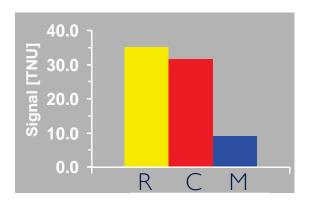
Borexino at Gran Sasso

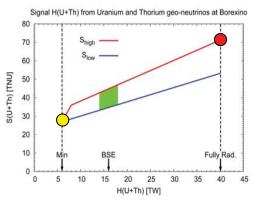


• A 300-ton liquid scintillator underground detector, running since may 2007.

- Signal, mainly generated from the crust, is comparable to reactor background.
- From BSE expect 5 7 events/year*
- In about two years should get 3σ evidence of geo-neutrinos.





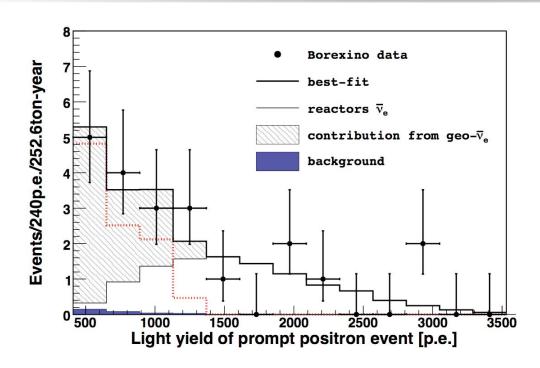


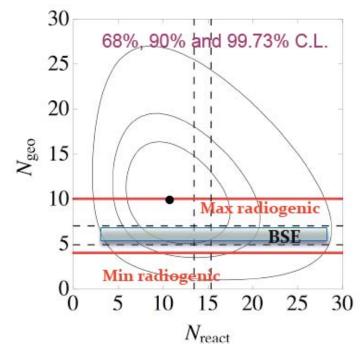






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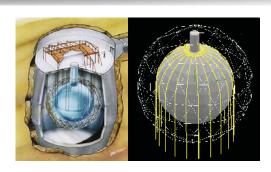


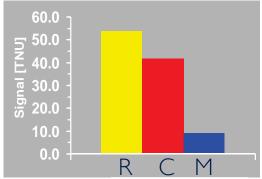
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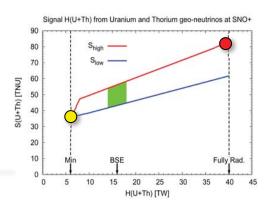


SNO+ at Sudbury

- A 1000-ton liquid scintillator underground detector, obtained by replacing D_2O in SNO.
- The SNO collaboration has planned to fill the detector with LS in 2009
- 80% of the signal comes from the continental crust.
- From BSE expect 28 38 events/year*
- It should be capable of measuring U+Th content of the crust.
- * assuming 80% eff. and 1 kTon CH₂ fiducial mass Chen, M. C., 2006, Earth Moon Planets 99, 221.











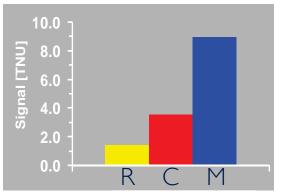


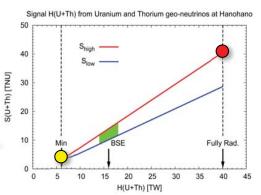
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Hanohano at Hawaii

- Project of a 10 kiloton movable deep-ocean LS detector
- ~ 70% of the signal comes from the mantle
- From BSE expect 60 100 events/year*
- It should be capable of measuring U+Th content of the mantle

The same of the sa









^{*} assuming 80% eff. and 10 kTon CH_2 fiducial mass

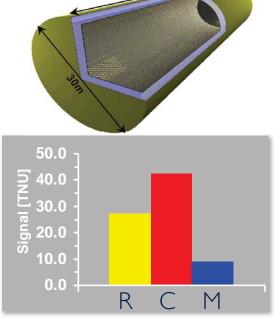
J. G. Learned et al. – ``XII-th International Workshop on Neutrino Telescope'', Venice, 2007

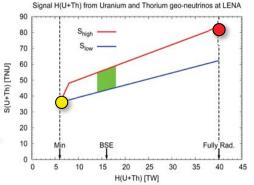


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LENA at Pyhasalmi

- Project of a 50 kiloton underground liquid scintillator detector in Finland
- 80% of the signal comes from the crust
- From BSE expect 800 1200 events/year*
- LS is loaded with 0.1% Gd which provides:
 - better neutron identification
 - moderate directional information





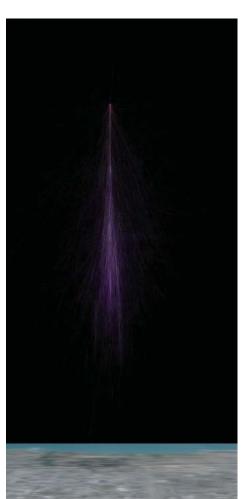
K.A. Hochmuth et al. - Astropart.Phys. 27 (2007) - arXiv:hep-ph/0509136; Teresa Marrodan @Taup 2007

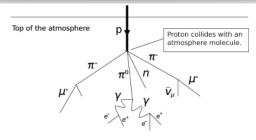


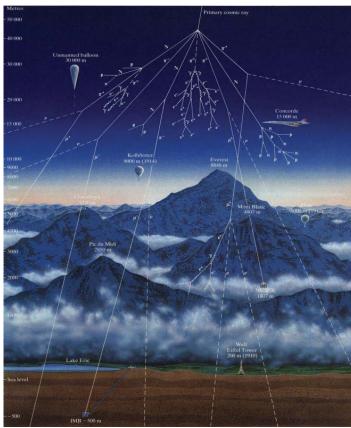


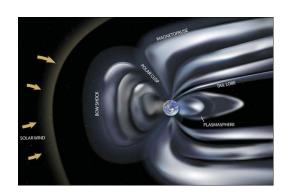
^{*} For 2.5 10³³ free protons and assuming 80% eff.







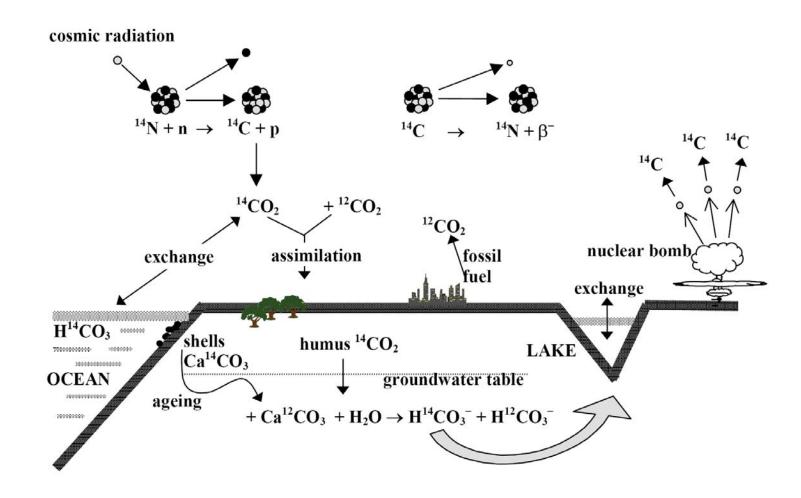








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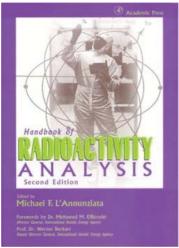


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COSMIC BACKGROUND REDUCTION IN THE RADIOCARBON MEASUREMENTS BY LIQUID SCINTILLATION SPECTROMETRY AT THE UNDERGROUND LABORATORY OF GRAN SASSO

Wolfango Plastino¹ • Lauri Kaihola² • Paolo Bartolomei³ • Francesco Bella⁴

ABSTRACT. Radiocarbon measurements by two 1220 QuantulusTM ultra low background liquid scintillation spectrometers were performed at the underground laboratory of Gran Sasso and the Radiocarbon Laboratory of E.N.E.A.-Bologna to study the efficiency and background variations related to measurement sites. The same configuration setup, i.e. the same center of gravity of the 14 C spectrum (SQP(I) = 410 ± 1) was obtained in both instruments. Many different background and modern standards with pure analytical benzene were used and spectra for 40 one-hour periods were obtained. The data indicates a background reduction of approximately 65% between the surface and underground laboratories, with no differences in the efficiency. Recording similar efficiencies in both spectrometers is probably due to fairly identical photomultiplier characteristics. The cosmic noise reduction observed at the laboratory of Gran Sasso makes it possible to perform high precision 14 C measurements and to extend for these idealized samples the present maximum dating limit from 58,000 BP to 62,000 BP (5 mL, 3 days counting).



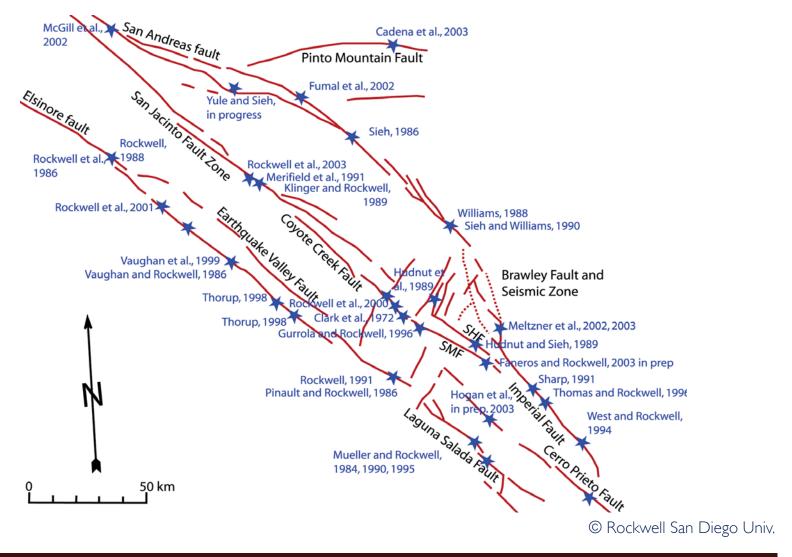
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Sample	Count rate (cpm)	Count error (cpm)	Modern activity (dpm)	Modern activity error (dpm)	Eff (%)	FM	fM	T _{max} (BP)
L1A H1A	0.278 12.949	0.022 0.148	8.853	0.088	80.93	25,531.503	17	48,200
L1A H1A	0.059 12.282	0.010 0.144	8.540	0.094	76.76	99,734.000	35	54,000
L3A H3A	0.398 39.140	0.026 0.257	27.068	0.161	83.28	17,419.104	43	55,900
L3A H3A	0.150 38.235	0.016 0.254	26.609	0.166	81.35	44,052.587	69	59,600
L5A H5A	0.655 65.206	0.033 0.332	45.101	0.209	83.60	10,676.865	57	58,000
L5A H5A	0.235 63.874	0.020 0.328	44.464	0.215	81.89	28,580.867	92	61,900





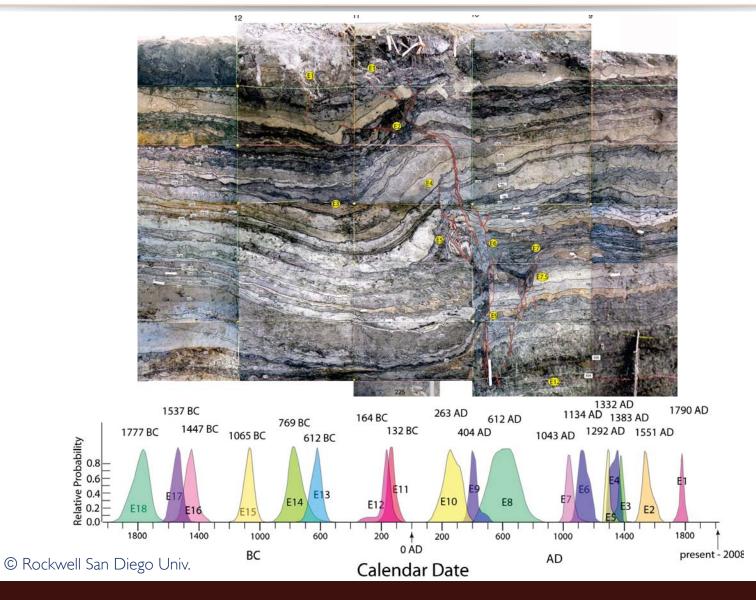








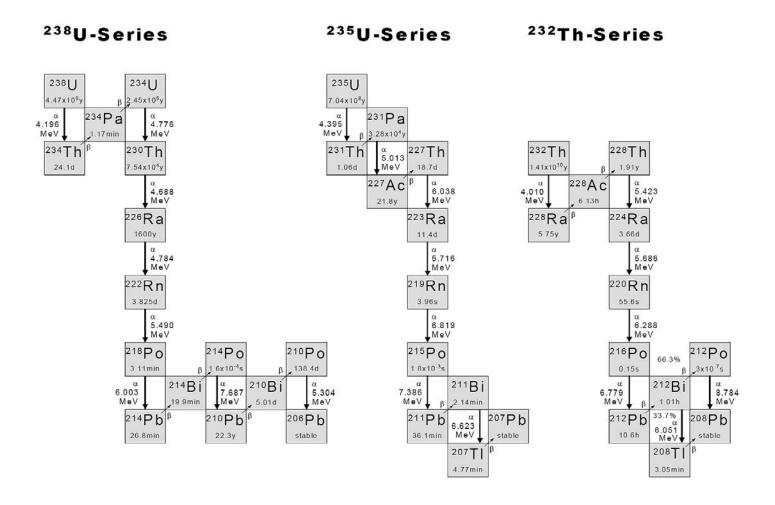








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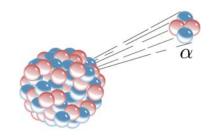
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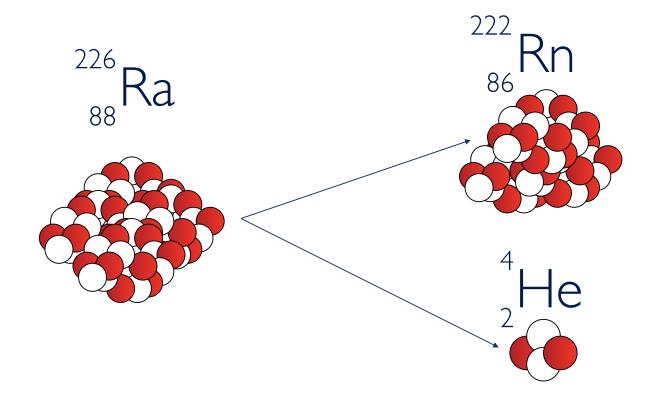




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Alpha Decay

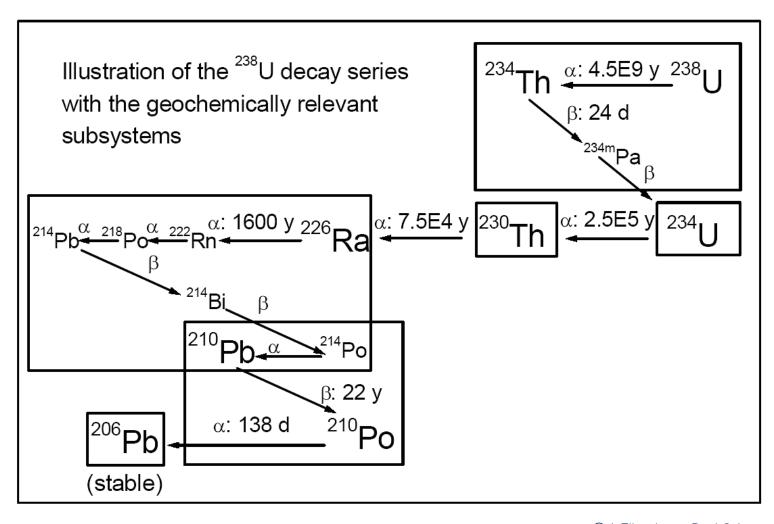




$$_{z}^{A}X \xrightarrow{A-4} Y + _{2}^{4}He$$



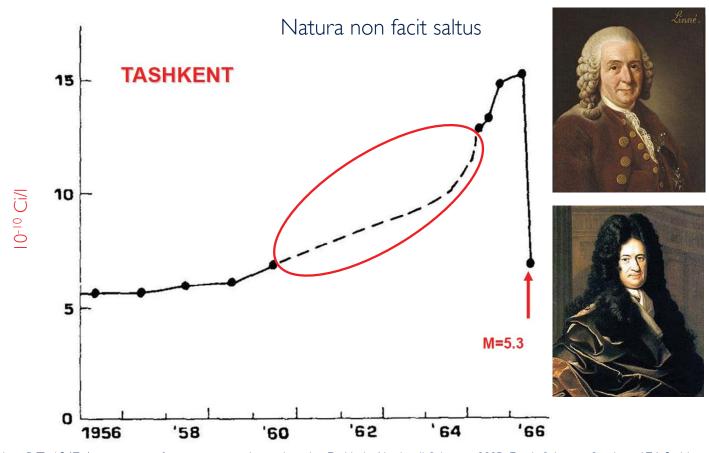




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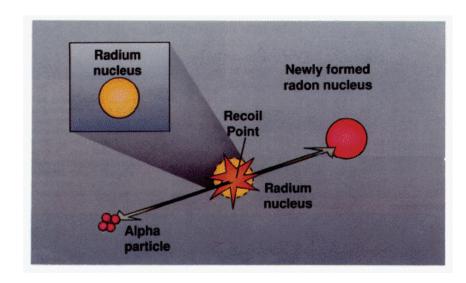


Ulomov, V.I., Mavashev, B.Z., 1967. A precursor of a strong tectonic earthquake. Doklady Akademii Sciences SSSR, Earth Sciences Sections 176, 9–11.



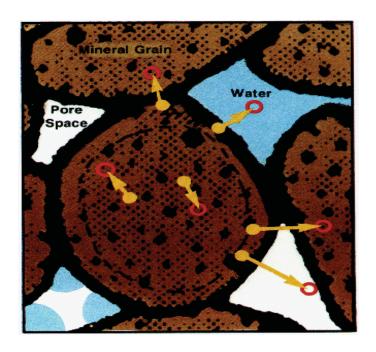


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The emanating power of rocks is defined as the ratio between the amount of radon escaping from the solid matrix and that produced by radioactive decay

Radium decay involves the release of the excess energy which is shared between the α particle which forms (98.1%), and the new radon atom



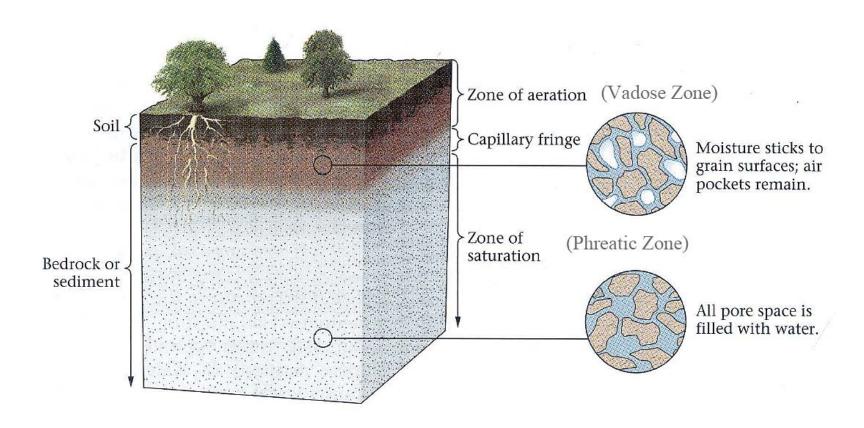
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Earth scienc

Radon and rock deformation

Evelyn Roeloffs

hat happens when stress is applied to rocks in the Earth's crust so that the crust deforms? This is a question tackled by Trique et al. on page 137 of this issue¹. They have used a natural laboratory in the French Alps — the Roselend reservoir — to monitor the geophysical signals that result from the greater or lesser pressure on the underlying crust exerted by the weight of water in the reservoir. This area is not itself prone to earthquakes. But the broader interest of this work is in what it may tell us about the events, induced by crustal deformation, that precede earthquakes.

The ability to predict earthquakes is of course highly desirable. But progress in this difficult and highly contentious science will depend on detecting and interpreting physical changes stemming from the processes

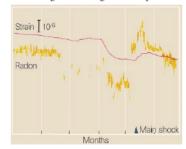


Figure 1 The radon and strain data for the magnitude-7 Izu–Oshima earthquake^{2,0} of 14 January 1978 show changes preceding the earthquake. But they do not match the model shown in Fig. 2; in particular, neither change is monotonic, and in both cases the pre-earthquake change exceeds that produced by the earthquake itself.

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of earthquake generation. Many possible precursors have been reported, but seismologists are sceptical of those that are not clearly linked to crustal deformation. This 'unproven' category includes the well-documented precursory decrease and increase of radon concentration before the 1978 Izu–Oshima earthquake in Japan² (Fig. 1), as well as the controversial assertion that

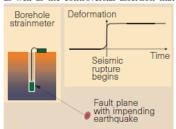


Figure 2 Rock friction, which depends on slip rate and sliding-induced changes on a fault surface, implies that seismic slip should be preceded by accelerating aseismic slip near the hypocentre of an impending earthquake. Sufficient aseismic slip would produce nearsurface deformation detectable by a borehole strainmeter. Compared with the strain step recorded at the time of the earthquake, the precursory strain signal would be in the same direction but of much smaller amplitude. A magnitude-5 earthquake, 10 km deep, produces maximum near-surface strain of about 10⁻⁷ at a site 5 km from its fault plane; strain increases 30-fold for each unit increase of magnitude, but falls off as the third power of distance from the source. Estimates of pre-seismic slip duration and amplitude range widely because frictional parameters of natural faults are poorly known.

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Seismologists expect earthquake precursors to take the form of transient crustalstrain signals from 'aseismic' fault slip near the earthquake's nucleation point (that is, fault slip that is too slow to radiate seismic waves) (Fig. 2). Numerical simulations show, however, that such signals would be exceedingly small4. Even the best existing instruments - borehole strainmeters with resolution exceeding a part per billion would need to be within a few kilometres of the impending earthquake's epicentre to detect this aseismic strain. Although strain changes preceding two California earthquakes have been identified^{5,6}, they don't resemble the expected signals.

Proponents of earthquake prediction maintain that changes in radon emission, or in electrical or magnetic fields, represent a natural amplification of pre-earthquake deformation under special geological conditions. For example, the conductance by rock fractures of water or gas is proportional to the third power of the fracture's aperture⁷. Fluid flow past ions adsorbed on rock surfaces produces an electric field, termed a 'streaming potential', that varies with pressure gradient and permeability8. Fluid, gas or electromagnetic measurements might thus detect deformation indirectly, albeit at localized sites and with amplitudes related nonlinearly to strain.

Silver and Wakita⁹ list many potential examples of such pre-earthquake 'strain indicators'. Unfortunately, these indicators are irreproducible: they can be detected only in certain locations, but in any one location earthquakes recur infrequently. What is needed is evidence that transient strain leads consistently, if not linearly or uniformly, to observable phenomena. The radon, electrical and ground-tilt measurements from Roselend lake constitute this kind of reproducible evidence.

The shallow crust's reaction to large changes in lake level may also illuminate the

NATURE VOL 399 13 MAY 1999 www.nature.com

...to be continued



