



**The Abdus Salam  
International Centre for Theoretical Physics**

The International Union of Geodesy and  
Geophysics



**2339-6**

**Workshop on Atmospheric Deposition: Processes and Environmental Impacts**

*21 - 25 May 2012*

**Land-Atmosphere Exchange of Trace Gases and Particles; Principles,  
Measurement and interpretation of fluxes**

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Gases and Particles; Principles,  
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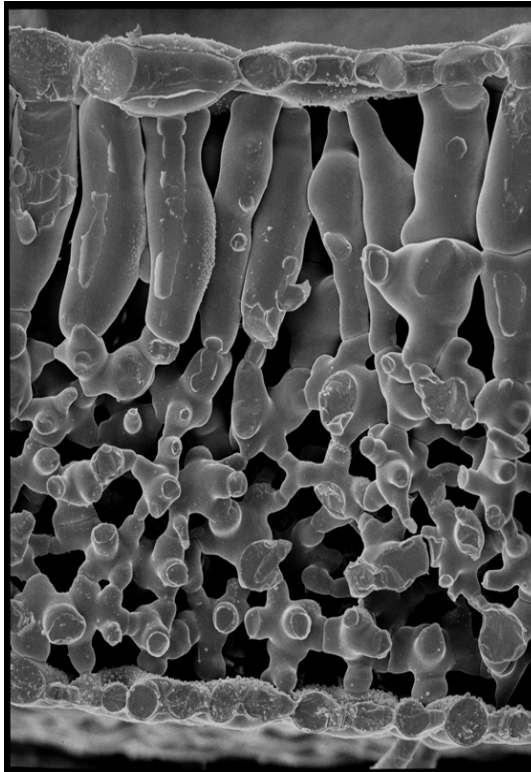
# Content

- **Introduction, the role of ecosystems in atmospheric composition**
- **Principles, the processes involved in surface-atmosphere exchange**
- **Resistances, deposition velocities and their components**
- **Measuring surface-atmosphere fluxes**
- **Ozone deposition to natural vegetation**
- **Sulphur deposition**
- **Surface-atmosphere exchange of ammonia (bi-directional exchange)**
- **Particle emission and deposition**
- **Methane and nitrous oxide exchange**

• **Introduction, the role of ecosystems in atmospheric composition**

- The earth is unique in having an atmosphere largely of biological origin
- For most trace gases exchange is regulated at the surface-atmosphere interface
- The policy needs have set priorities for research (SO<sub>2</sub>, NO<sub>2</sub>, O<sub>3</sub>.....)
- Surface –atmosphere exchange is a key source or sink of gases and particles contributing to Acidification eutrophication, ozone and other photochemical oxidants, climate change

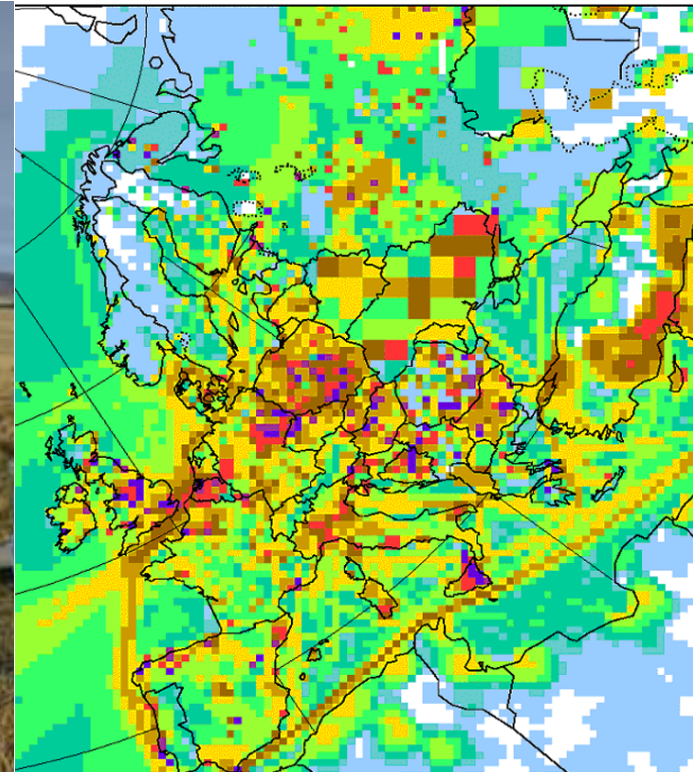
# SCALE



Molecular to  $\mu\text{m}$   
Controls the exchange



Measurement scale  
Landscape (100-1000m)

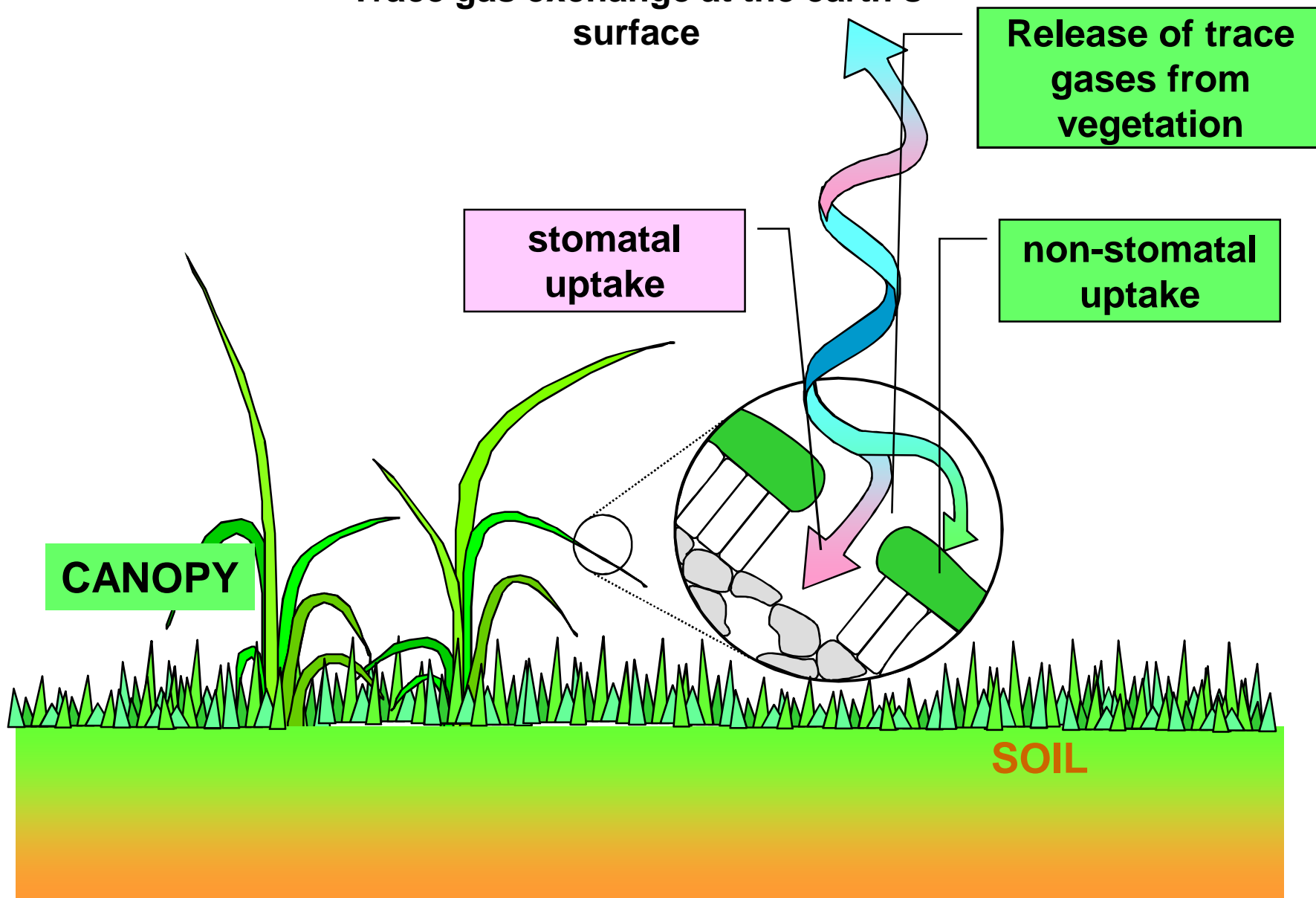


Application: region to  
continent 100km to 10000km

# The transfer process

- Turbulent transfer in the free atmosphere
- For trace gases, molecular diffusion becomes important very close to surfaces (<mm)
- At the surface for many of the trace gases chemical interaction with the surface components and water are very important

# Trace gas exchange at the earth's surface

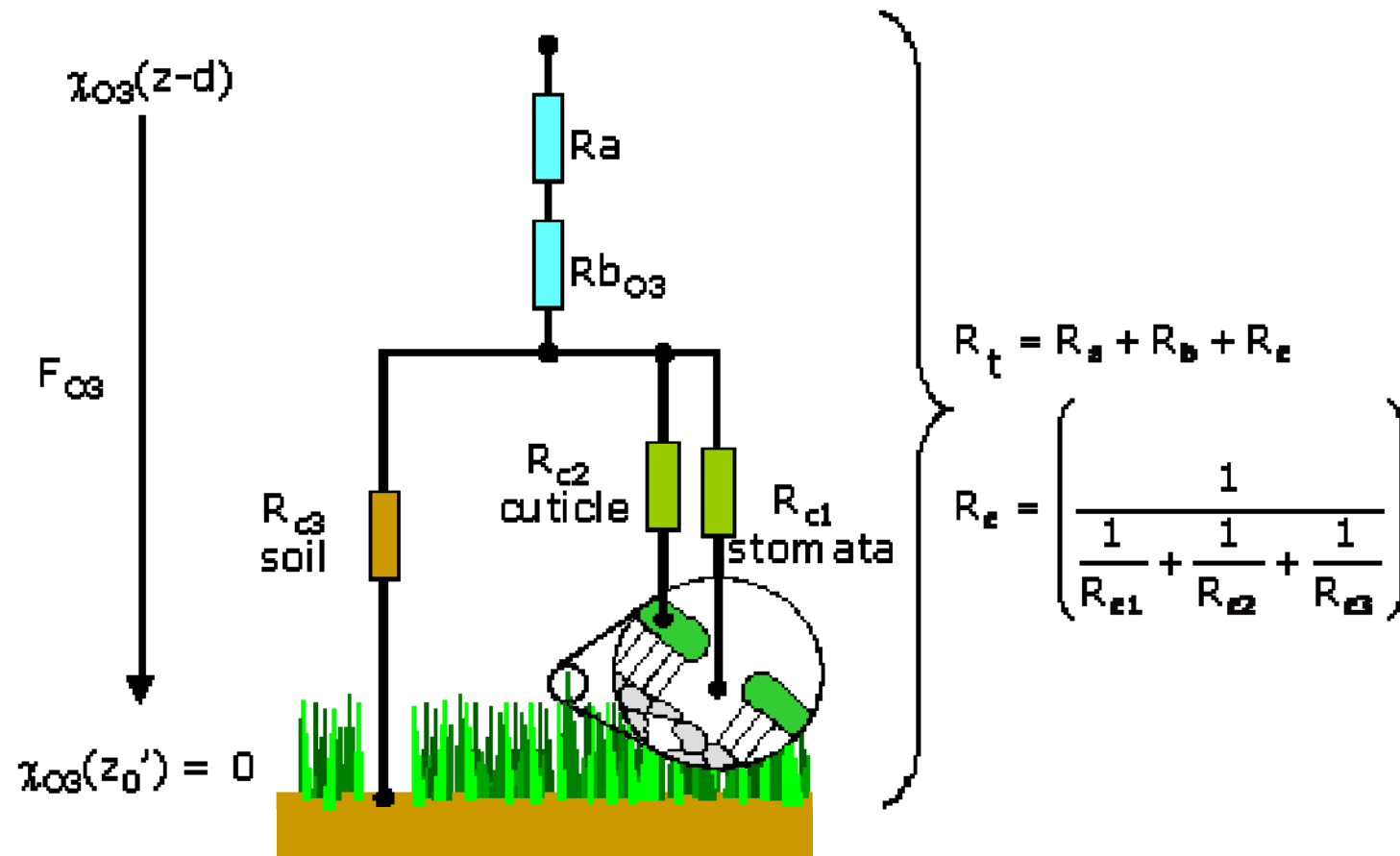


# Dry Deposition

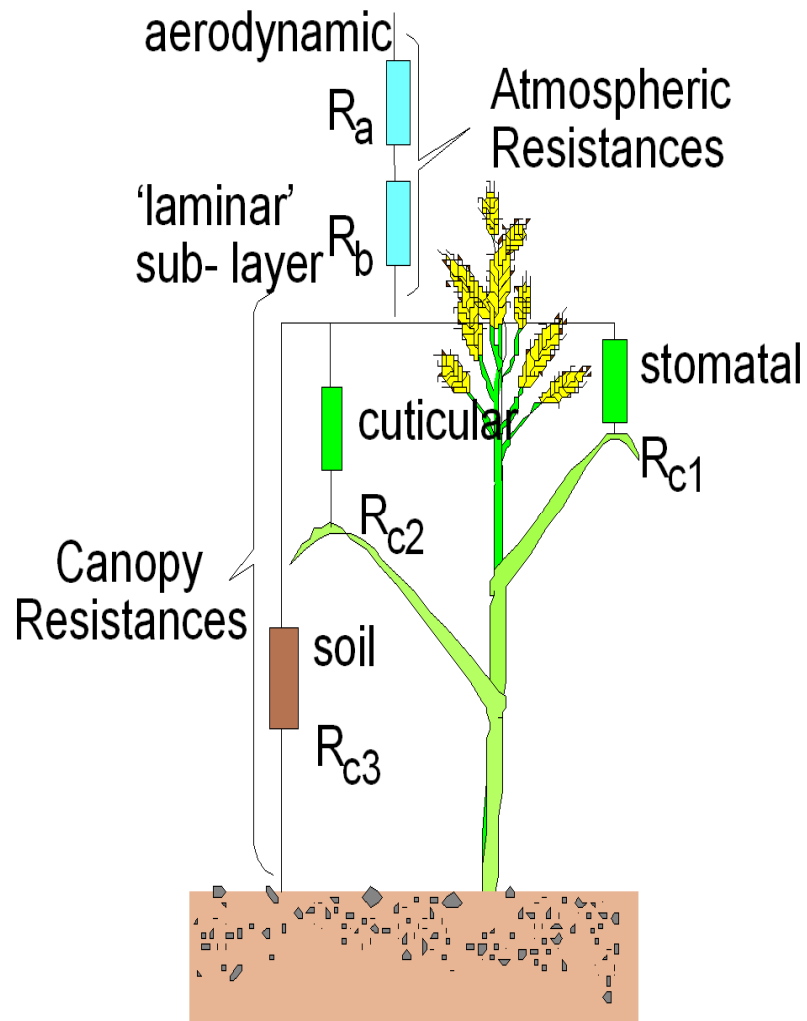
Definition: The direct transfer of pollutants to terrestrial surfaces



# Simple resistance models



# The resistance analogy



$$V_d = \frac{\text{flux}}{\text{concentration}} = \frac{1}{R_t}$$

$$R_t = R_a + R_b + (1/R_{c1} + 1/R_{c2} + 1/R_{c3})^{-1}$$



Contents lists available at ScienceDirect

## Atmospheric Environment

journal homepage: www.elsevier.com/locate/atmosenv



position interactions

## Review

## Atmospheric composition change: Ecosystems–Atmosphere interactions

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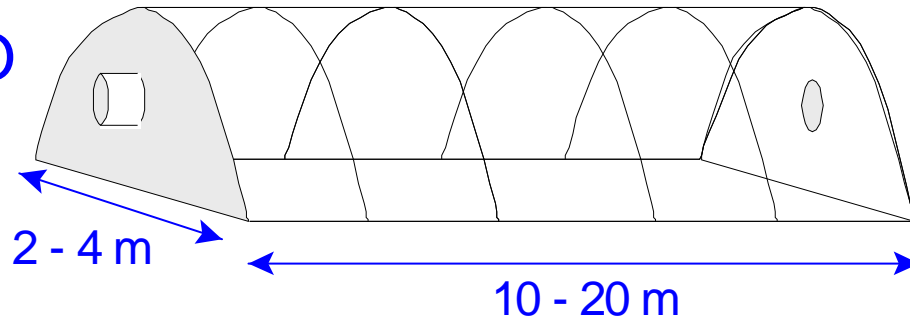
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# Flux measurement techniques

- Chamber methods    plot scale
- Micrometeorology....(field to landscape)
- Boundary layer budget (landscape to region)
- Remote sensing (global)

- ② overcomes small scale variability
- ② provides substantial spatial averaging

Mega-chamber  $N_2O$   
20 - 80  $m^2$



CLIMEX -  
catchment "greenhouse"  
study  $NO$ ,  $NO_2$ ,  $(CH_3)$ ,  $CO_2$ ,  $O_3$



450  $m^2$



# Chamber Effects

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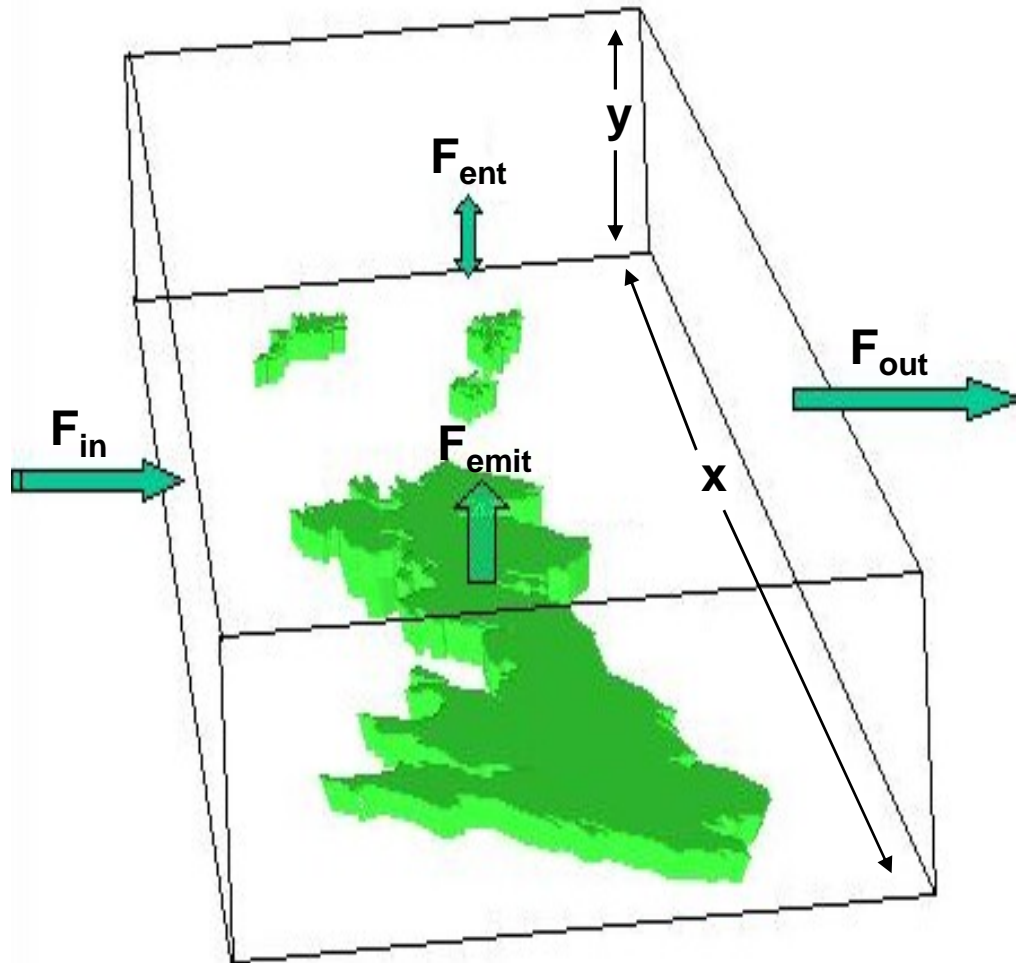
1.  $\Delta$ temperature +10 °C over 30 minutes

CH<sub>4</sub> emission varies with temperature but the chamber temperature and water temperature do not have the same response time so that long term data cannot be used to quantify this effect (< 10 % ?)

2.  $\Delta$ pressure

A very small effect as there is a large hole in the chamber wall to maintain pressure.

# Estimating Annual Emissions – Box Technique



- $F_{\text{emit}} = F_{\text{out}} - F_{\text{in}} - F_{\text{ent}}$
- $F_{\text{in}} = \iint u_{\text{in}}(y,z)c(y,z)dydz$
- $F_{\text{out}} = \iint u_{\text{out}}(y,z)c(y,z)dydz$
- $F_{\text{ent}} = 0$

$F_{\text{emit}}$  = Flux emitted

$F_{\text{in}}$  = Flux in

$F_{\text{out}}$  = Flux out

$u$  = wind speed

$c$  = concentration

# Atmospheric Mass Budgets at Regional Scales

## UPWIND

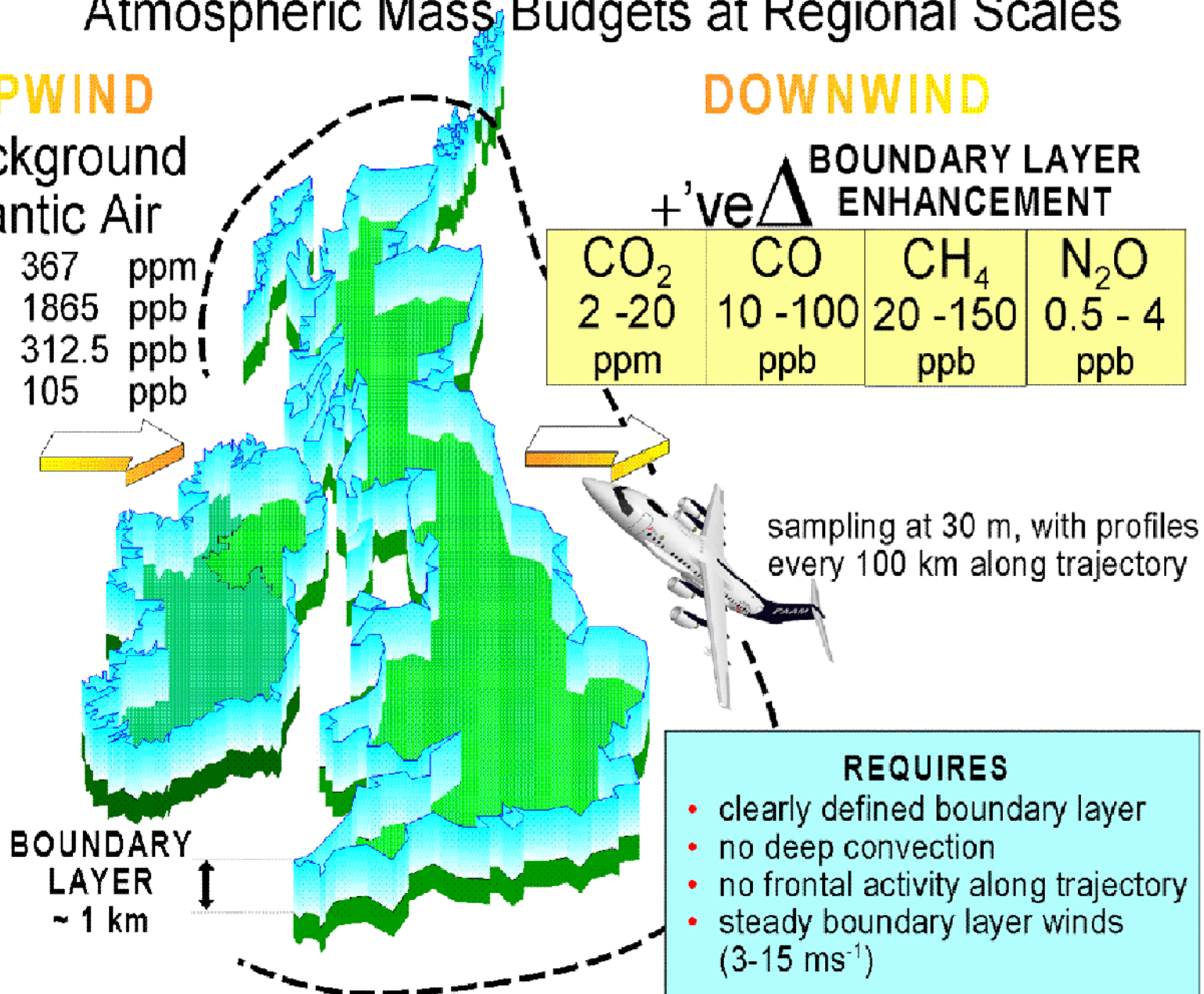
Background  
Atlantic Air

CO <sub>2</sub>	367	ppm
CH <sub>4</sub>	1865	ppb
N <sub>2</sub> O	312.5	ppb
CO	105	ppb

## DOWNWIND

+ve  $\Delta$  BOUNDARY LAYER  
ENHANCEMENT

CO <sub>2</sub>	CO	CH <sub>4</sub>	N <sub>2</sub> O
2 -20	10 -100	20 -150	0.5 - 4
ppm	ppb	ppb	ppb



BOUNDARY  
LAYER  
~ 1 km

sampling at 30 m, with profiles  
every 100 km along trajectory

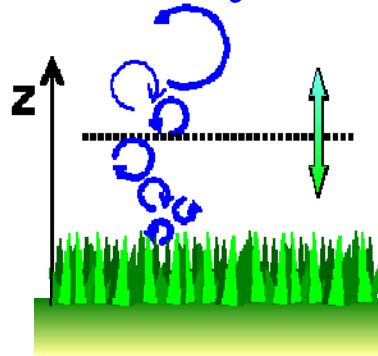
### REQUIRES

- clearly defined boundary layer
- no deep convection
- no frontal activity along trajectory
- steady boundary layer winds (3-15 ms<sup>-1</sup>)



# Micrometeorological Methods

## Eddy-covariance (EC)

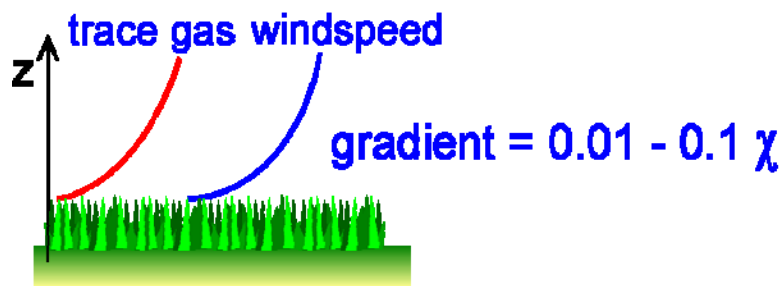


$$\overline{\chi'} = 0.01 - 0.05 \overline{\chi}$$

$$\text{flux}_{\chi} = \overline{w' \chi'}$$

$w'$  - fluctuation in vertical wind speed  
 $\chi'$  - deviation from mean concentration

## Aerodynamic flux-gradient



$$\text{flux}_{\chi} = K_{\chi} \frac{\partial \chi}{\partial z}$$

with stability correction:

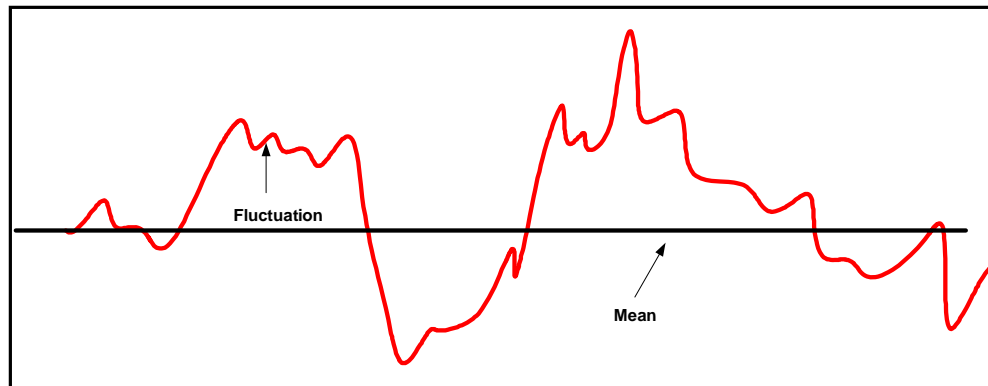
$$\text{flux}_{\chi} = k u_* \frac{\partial \chi}{\partial [\ln(z-d) - \Psi_H\{\zeta\}]}$$

# Eddy Covariance

- Pros
  - Direct , not inferred (e.g. K-theory), measure of the flux density between the atmosphere and biosphere
  - *In situ*
  - Quasi-continuous
  - Integrative of broad area
  - Introduces no artifacts, like chambers
- Con
  - Expensive
  - Suffers from errors in complex terrain or advective conditions
  - Modern chemical sensors require electrical power, limiting where the method can be employed
  - Not amenable for all trace gases, if fast sensor is not available

# Eddy Covariance Technique

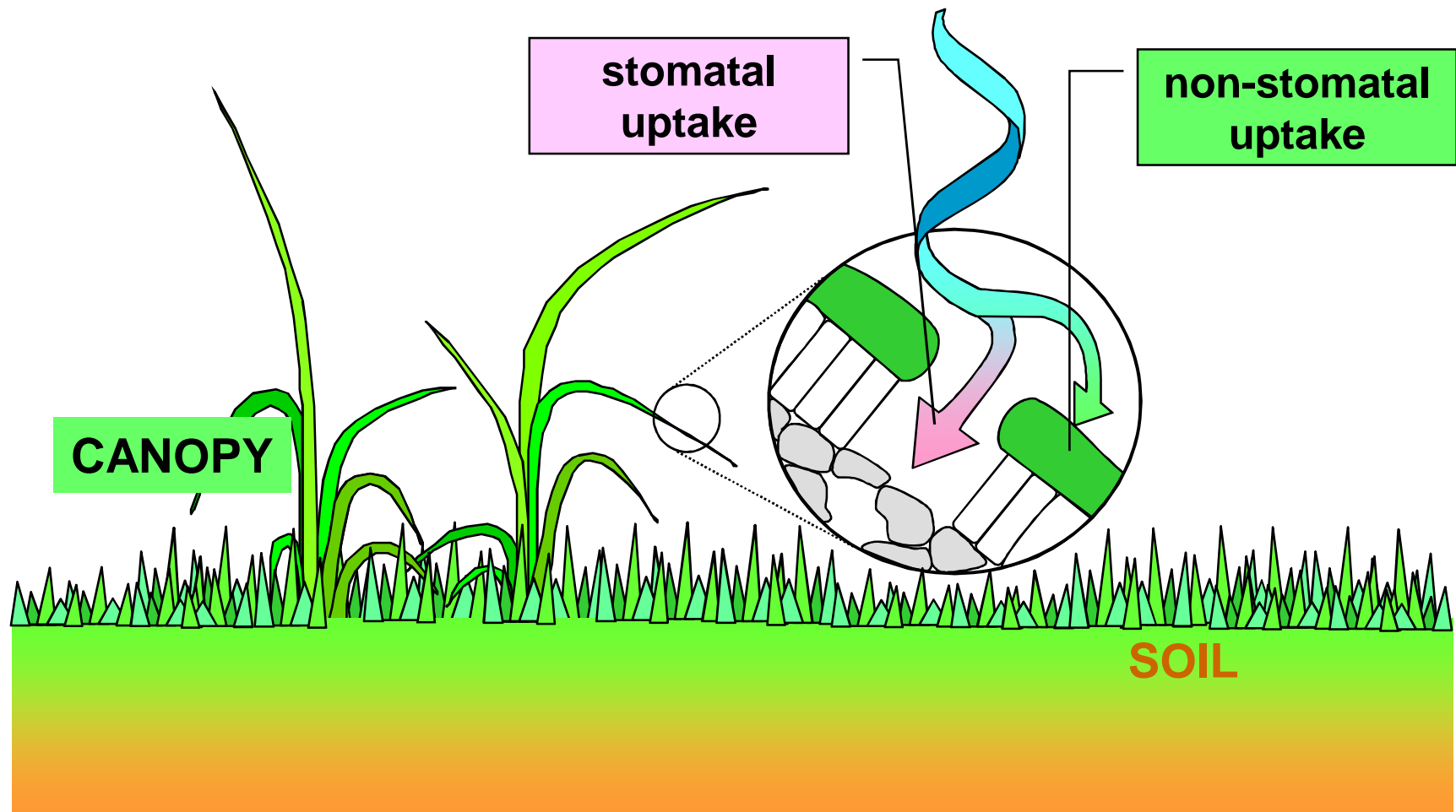
$$F = \overline{\rho w s} \sim \overline{\rho_a} \cdot \overline{w' s'}$$
$$s = \left( \frac{\rho_c}{\rho_a} \right)$$



# Ozone

- Identifying the different sinks for ozone at the surface
- Separating stomatal and non-stomatal fluxes to vegetation
- Understanding the effects of soil moisture, leaf temperature and humidity on ozone flux.
- Understanding non-stomatal ozone deposition
- The importance of O<sub>3</sub> deposition to vegetation in episode conditions and feedbacks from hot dry summers (2003).
- Long term measurements of ozone flux

## Ozone exchange at the earth's surface



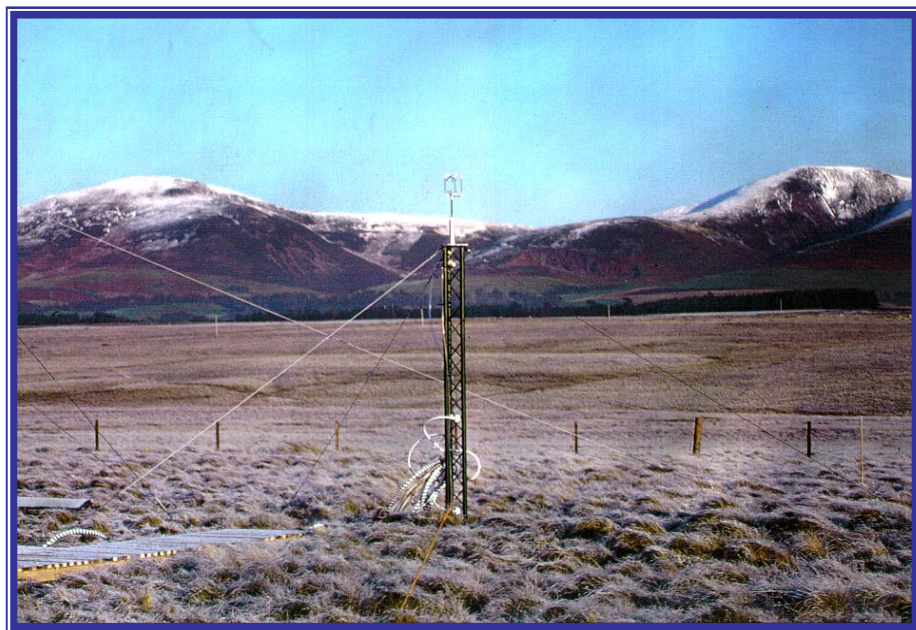
## ROFI: A New “Rapid Ozone Flux Instrument”

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- Measure  $O_3$  and  $H_2O$  fluxes using the eddy-correlation method.
- Low power instruments can be run using a wind turbine and solar panel.
- Allow measurement of total, stomatal and non-stomatal ozone deposition in remote locations without access to mains power.

# Auchencorth Moss



55° 47' 30" N, 3° 14' 20" W

Altitude 270 m a.s.l

Ombrotrophic mire, > 1000 ha

- moorland canopy height 0.2 – 0.4 m
- LAI 4-6 summer, 2-3 winter
- *Deschampsia flexuosa*, *molinia caerulea*,  
*eriphorum vaginatum*, *festuca ovina*,  
*calluna vulgaris*, *vaccinium myrtillus*,  
*carex nigra*, *sphagnum* spp.

❖ Continuous flux measurements from 1995 to 1998.

❖ Eddy-covariance (EC)  $u^*$ , C,  $\lambda E$ ,  $CO_2$

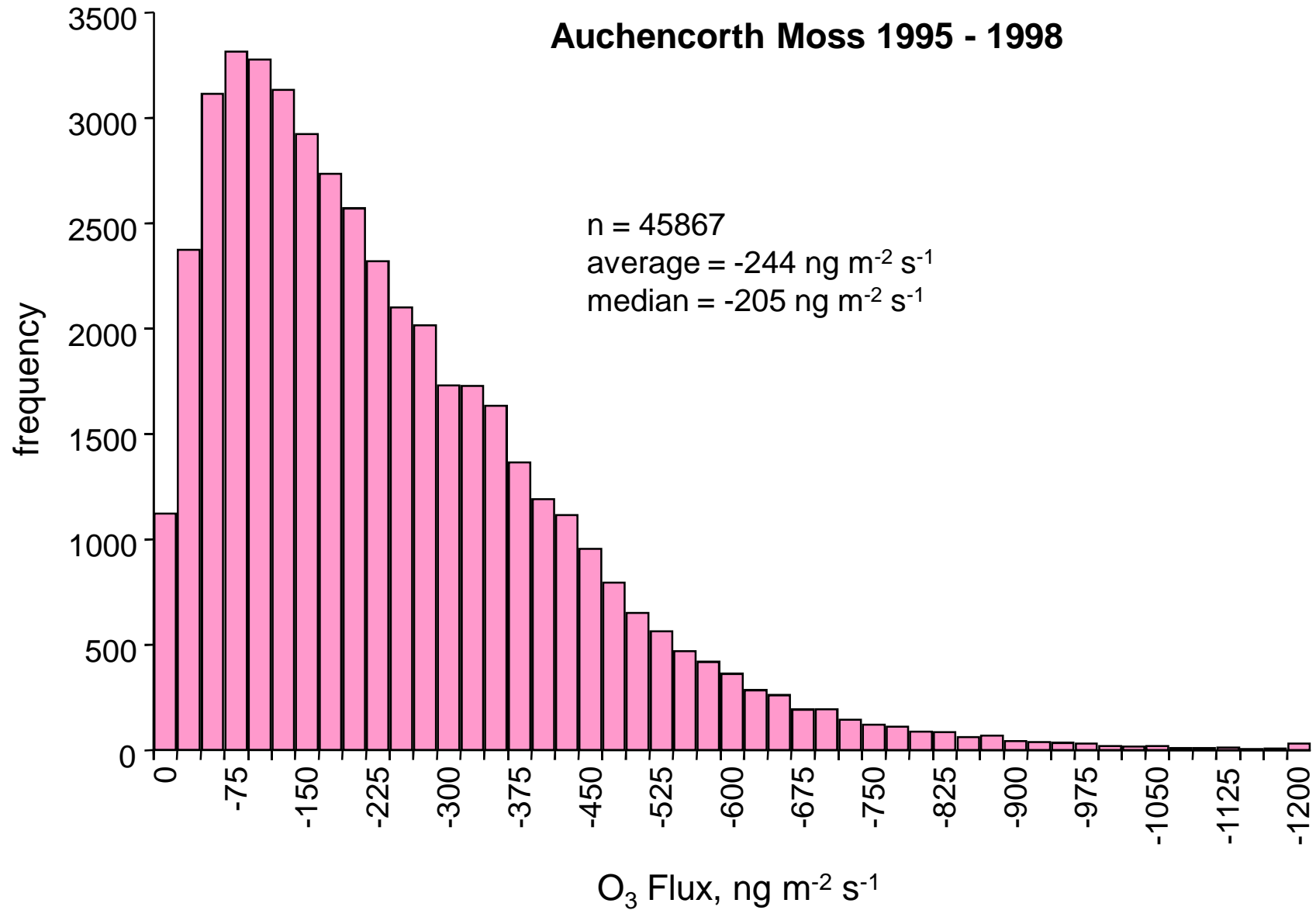
❖ EC & flux-gradient  $O_3$  (NO,  $NO_2$ ,  
 $SO_2$ ,  $NH_3$ )

❖ Bowen ratio  $\lambda E$ , C

Data capture, ~24000 hours of flux data:

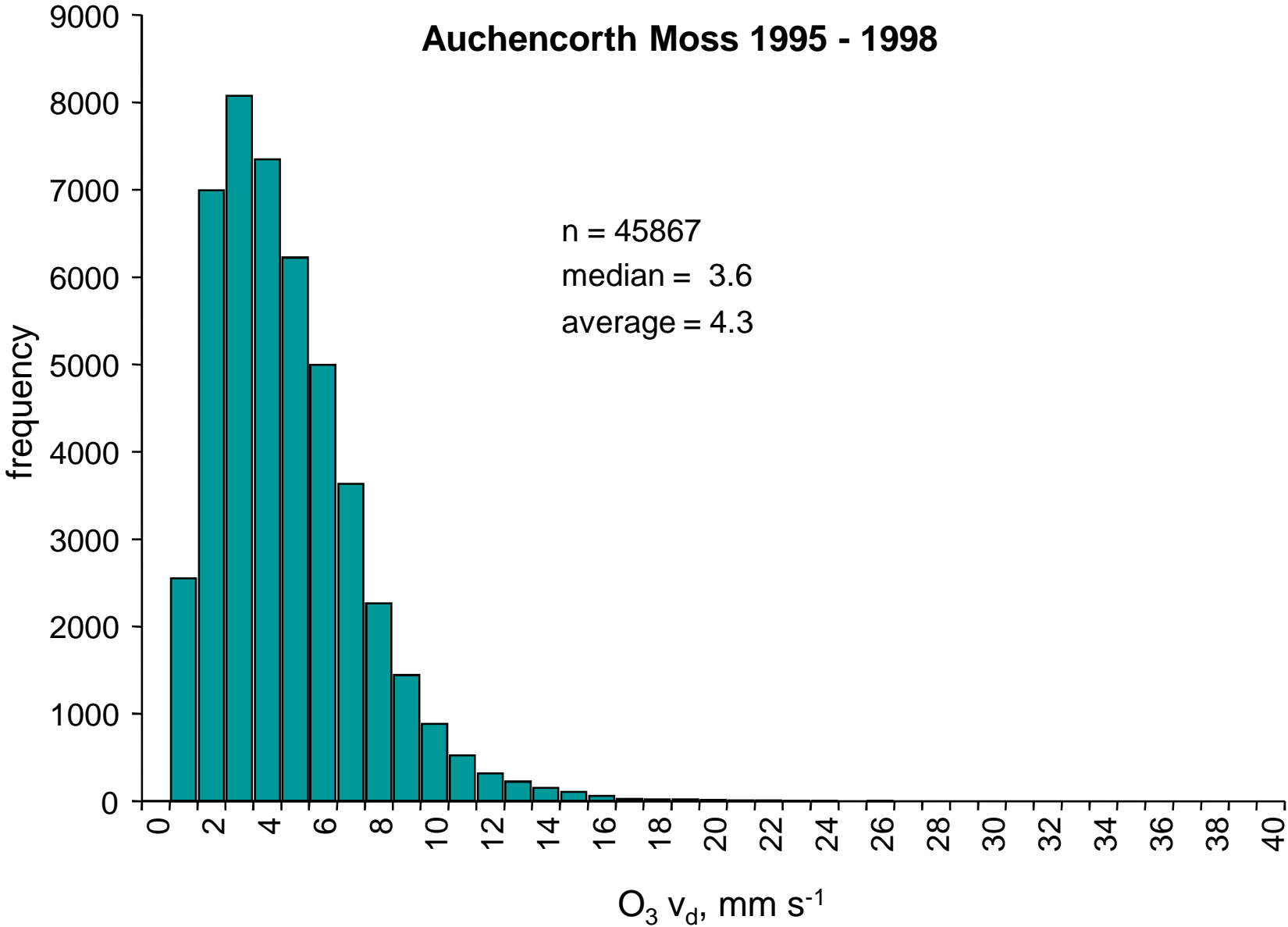
concentration	fluxes
94%	68%

# Frequency Distribution: O<sub>3</sub> deposition flux.

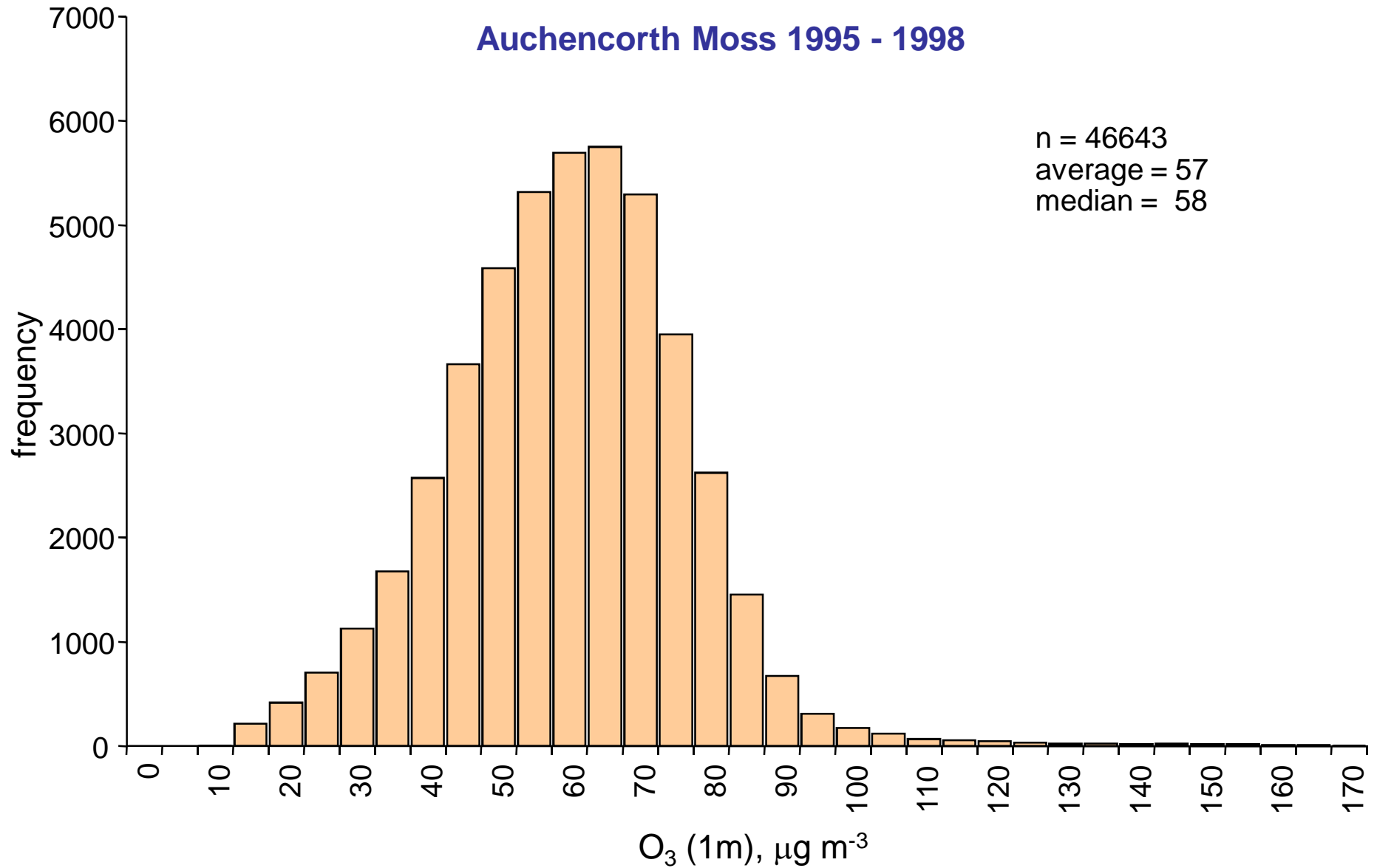




# Frequency Distribution: O<sub>3</sub> deposition velocity.

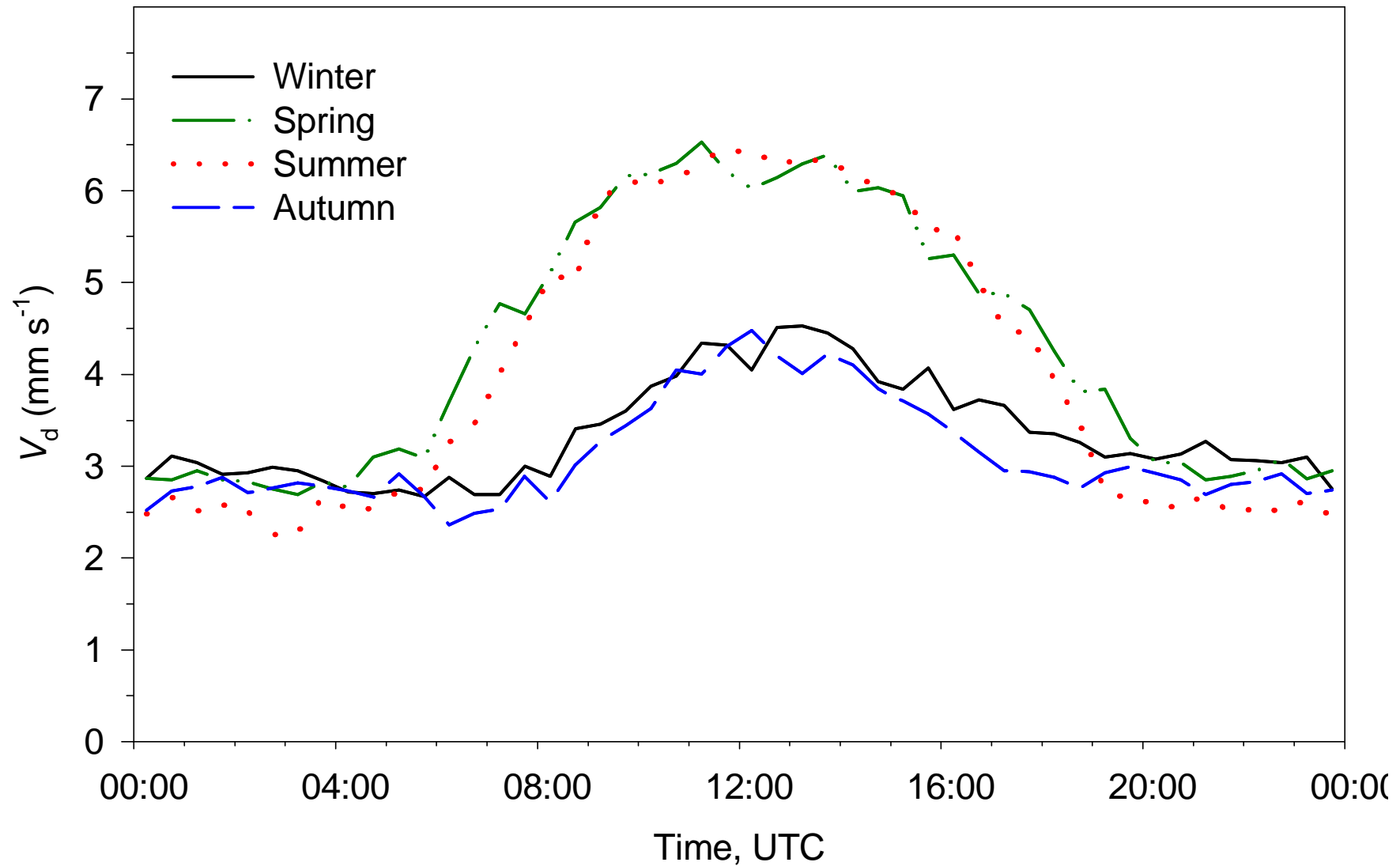


# Frequency Distribution: O<sub>3</sub> concentration.



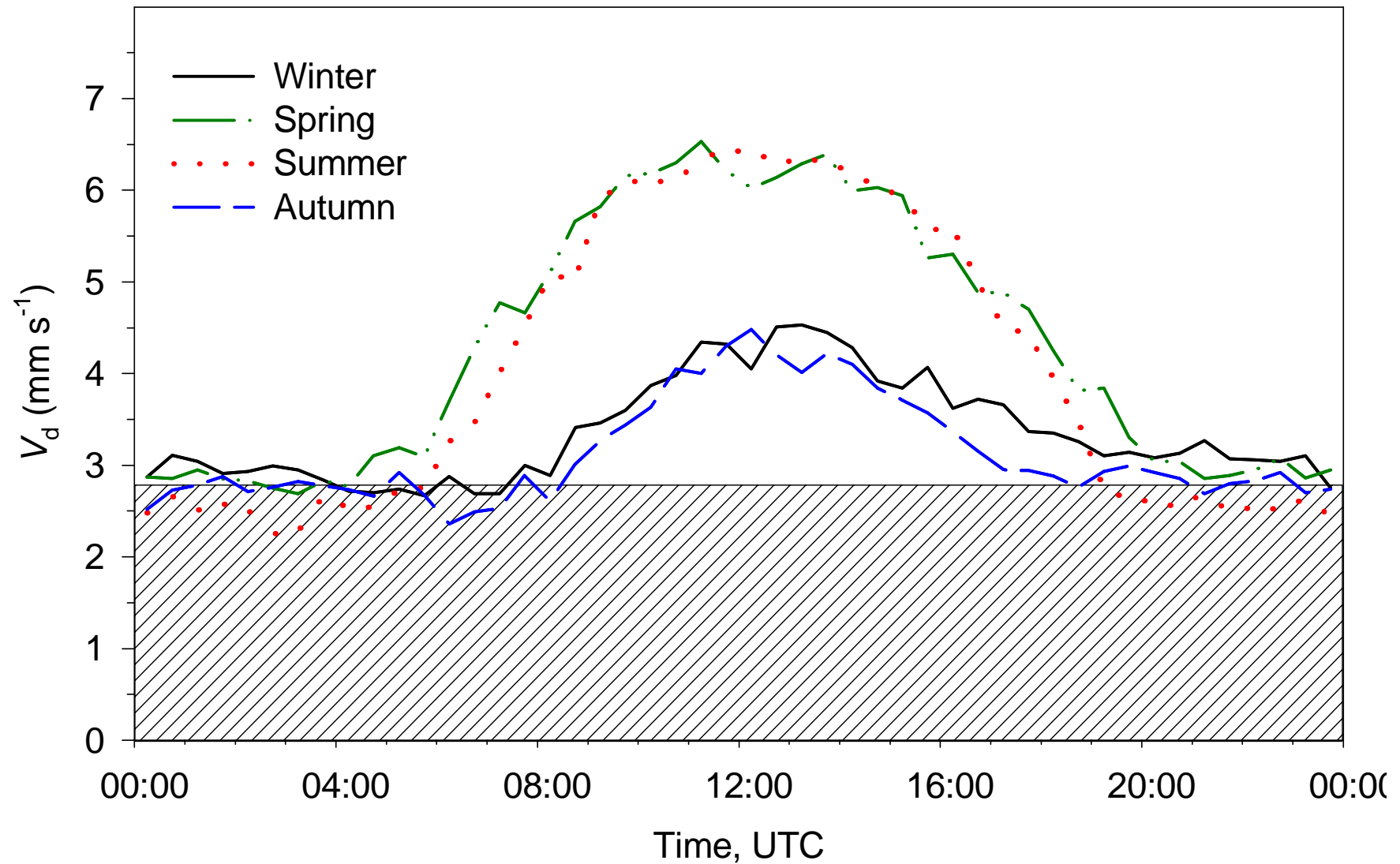
# 1995 – 98 Seasonal average deposition velocities at Auchencorth Moss

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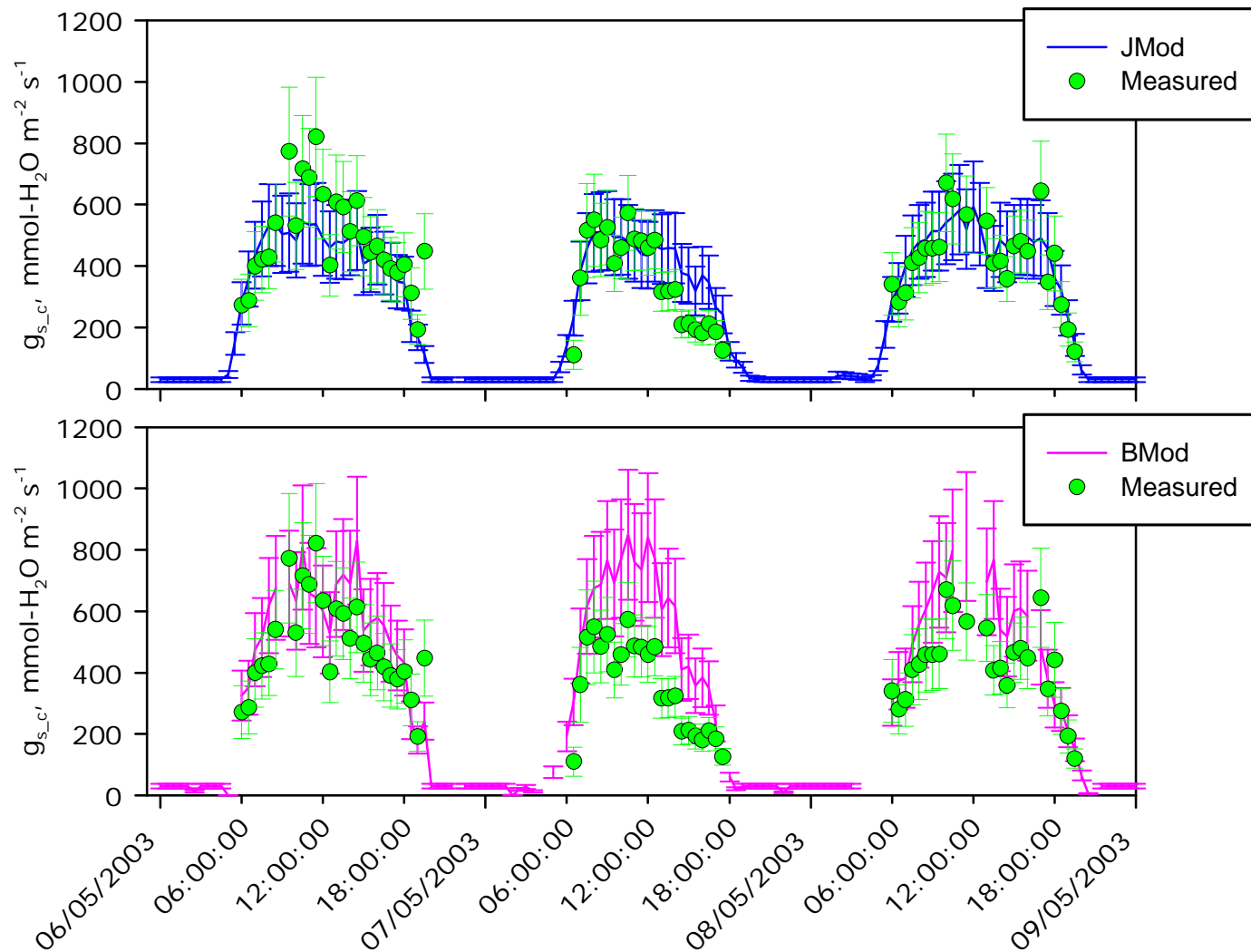


# 1995 – 98 Seasonal average deposition velocities at Auchencorth Moss

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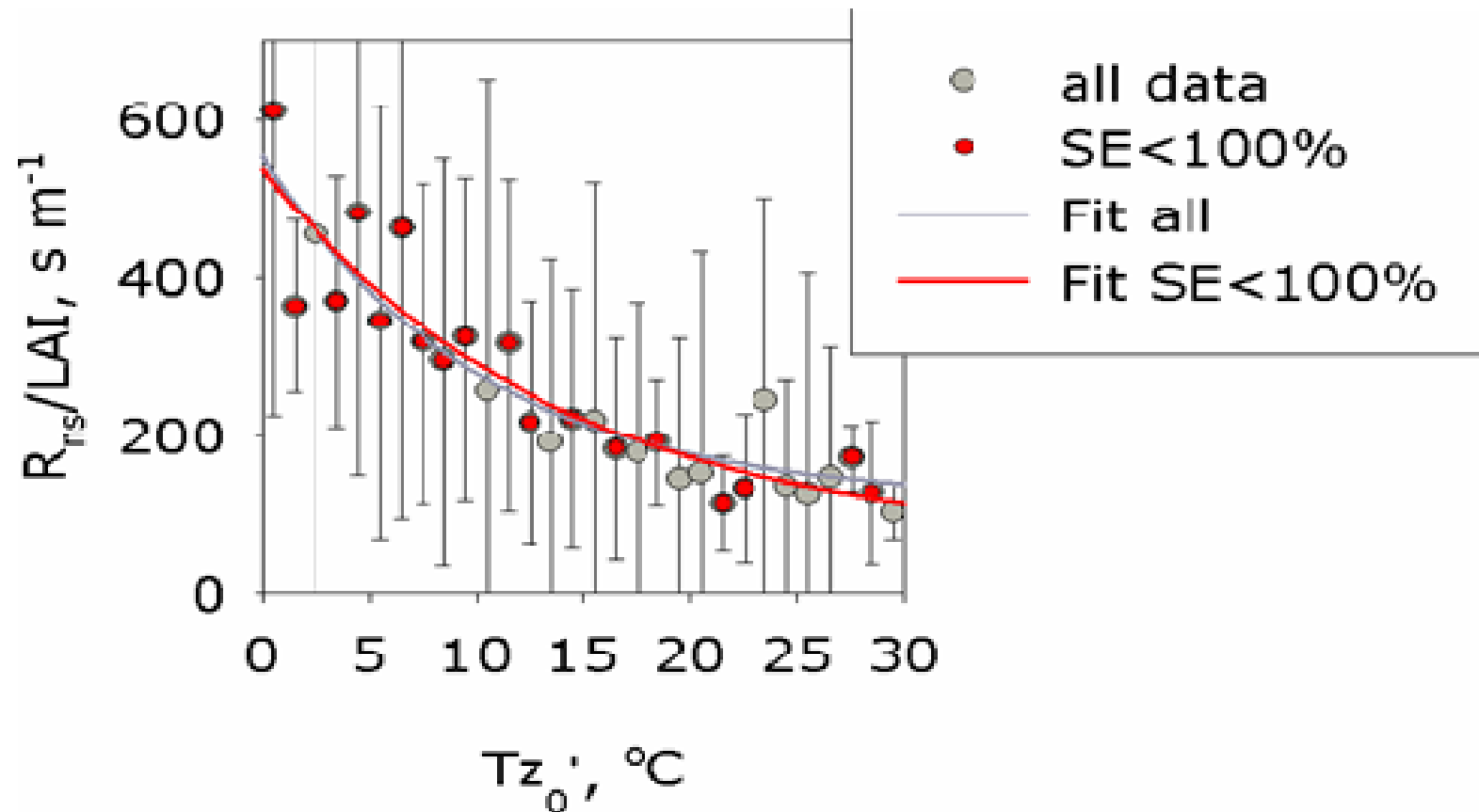


# Modelled and measured stomatal fluxes



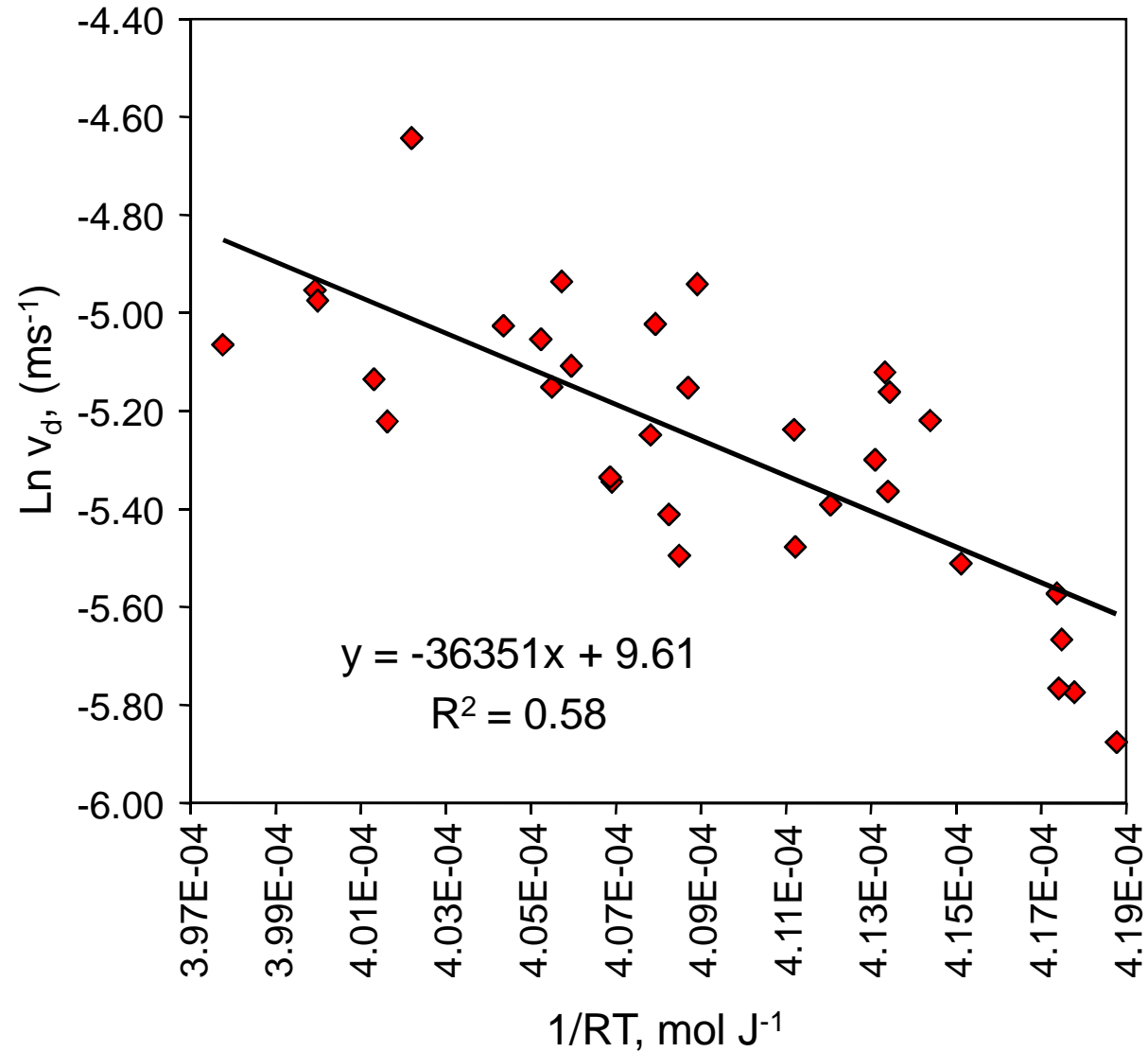
- The mechanisms of ozone depletion on leaf surfaces and within vegetation

## Effects of leaf surface temperature on canopy resistance



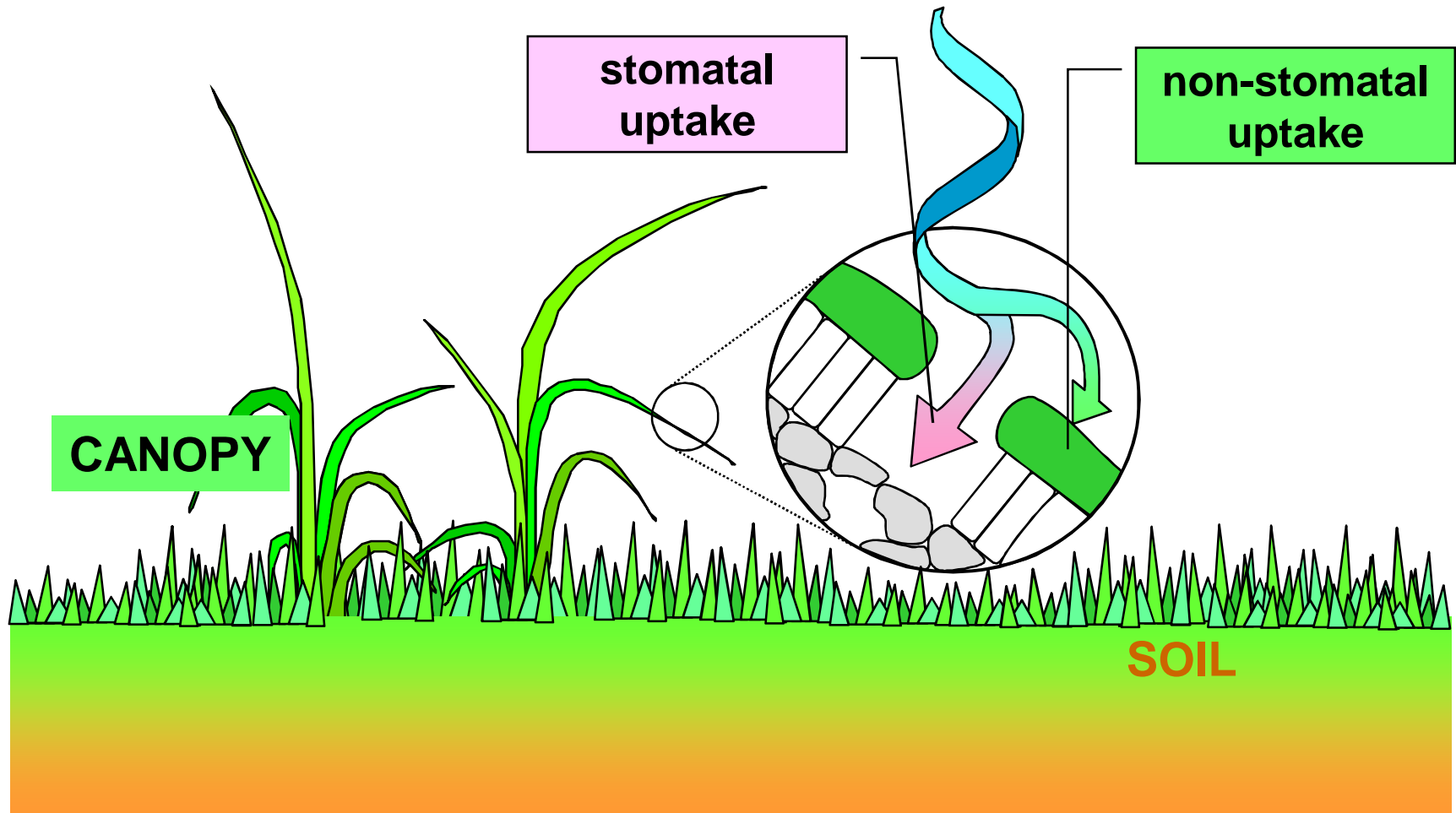
## Arrhenius plot for the reaction rate of ozone on non-stomatal surfaces

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# Pathways of Ozone Deposition

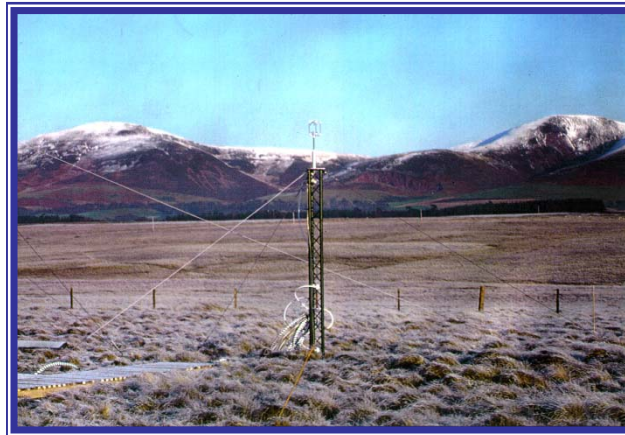
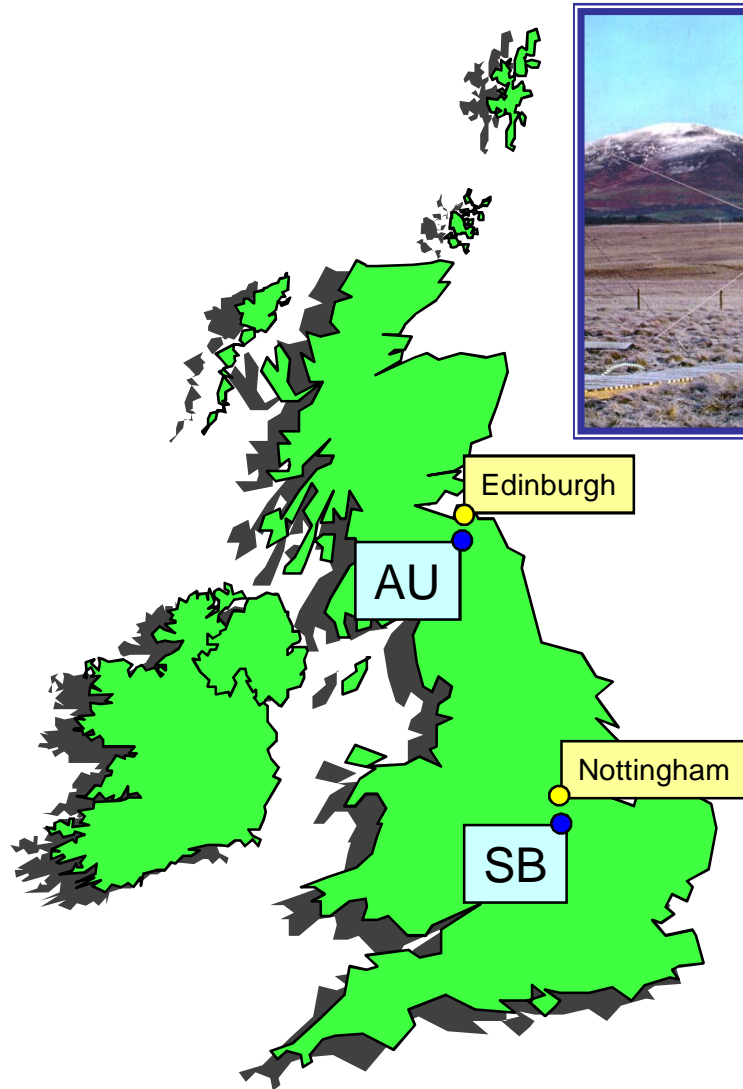


- Stomatal flux cf AOT40
- Ozone risk assessments by the two methods over Europe

# Sulphur Dioxide

- Processes regulating exchange at the surface, competition between stomata and leaf surface sinks, the effect of leaf surface wetness and the control of canopy resistance by chemistry in moisture on leaf surfaces.
- Non-linearities and the effects of  $\text{NH}_3/\text{SO}_2$  ratios at regional scales.
- Long-term trends in deposition velocity over Europe

# Auchencorth Moss (AU) & Sutton Bonnington (SB)



## Auchencorth Moss

55° 47' 30" N, 3° 14' 20" W

Altitude 270 m a.s.l

Ombrotrophic mire, > 1000  
ha



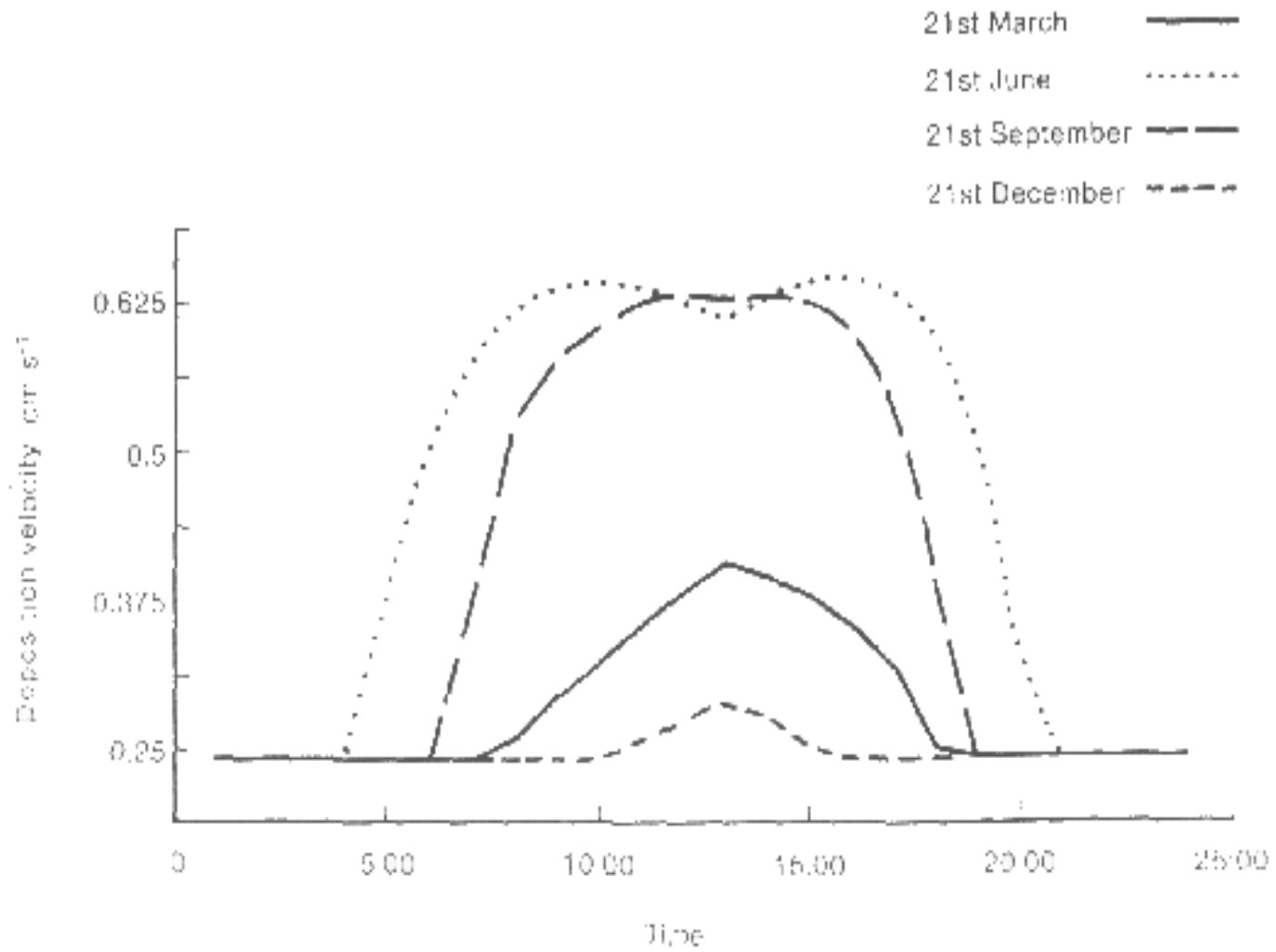
## Sutton Bonnington

52° 50' 24" N, 1° 15' E

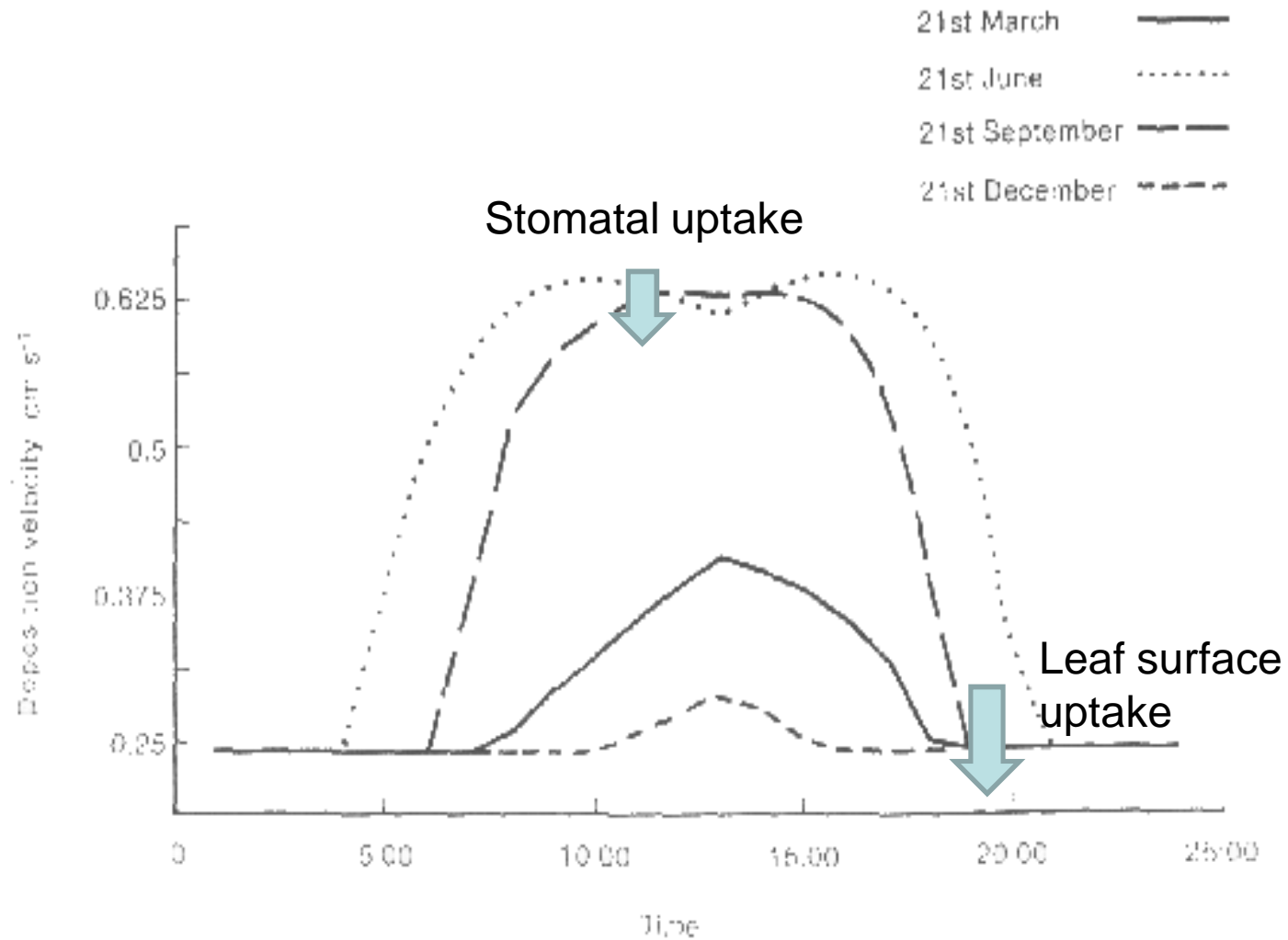
Altitude 45 m a.s.l

Arable farmland, > 4 ha

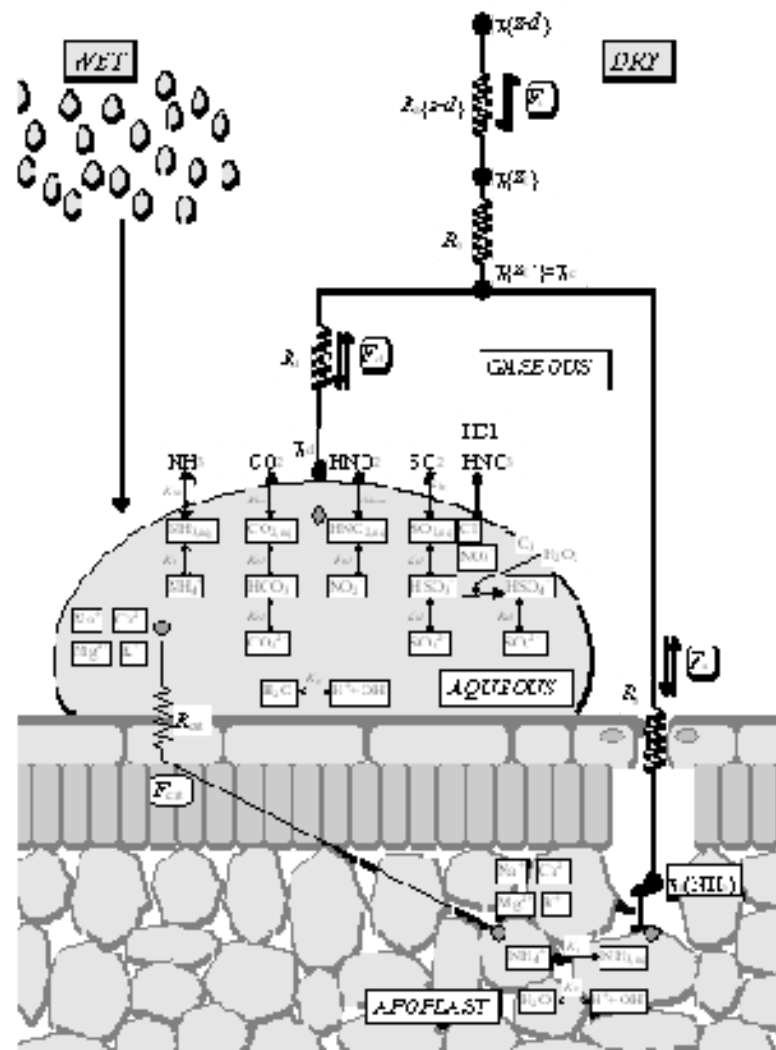
# SO<sub>2</sub> deposition on vegetation



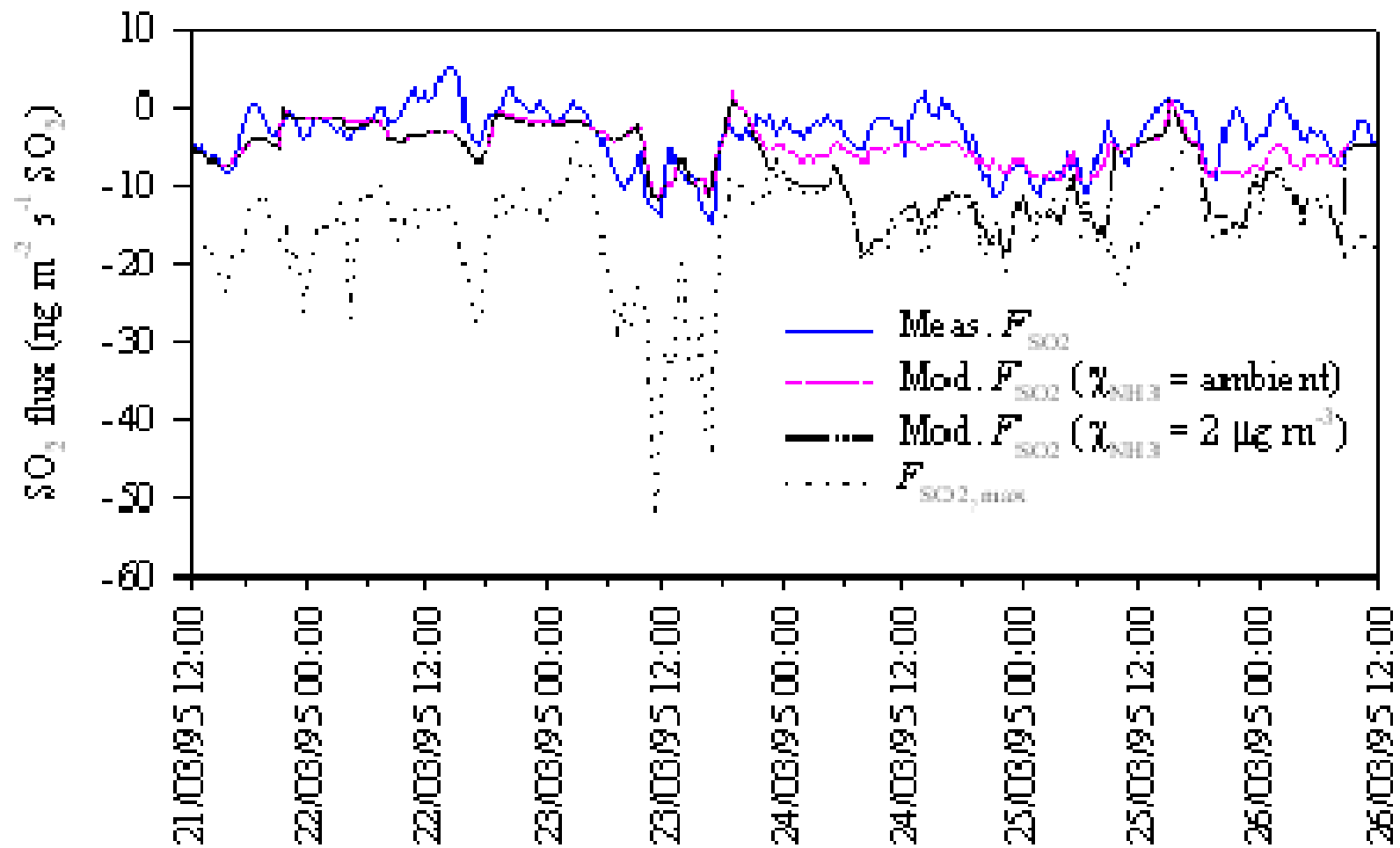
# SO<sub>2</sub> deposition on vegetation



# Leaf surface chemistry model for $\text{SO}_2$ deposition



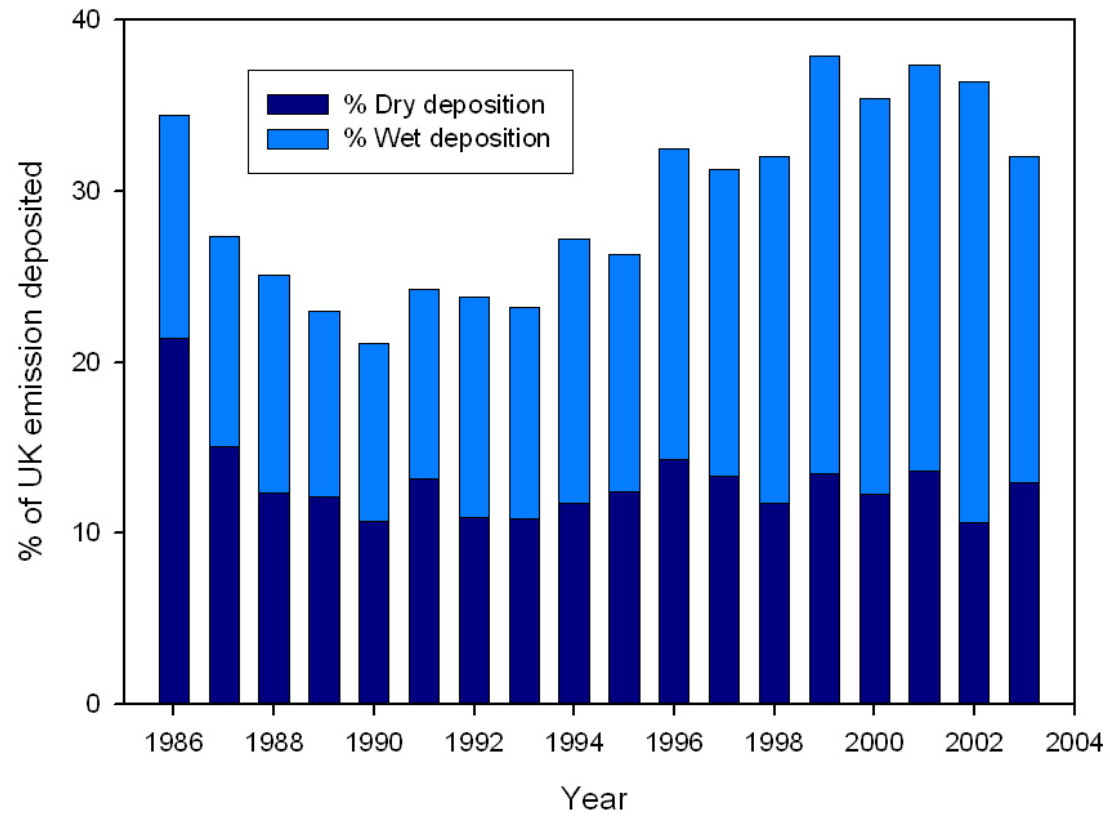
# Model- measurement comparison





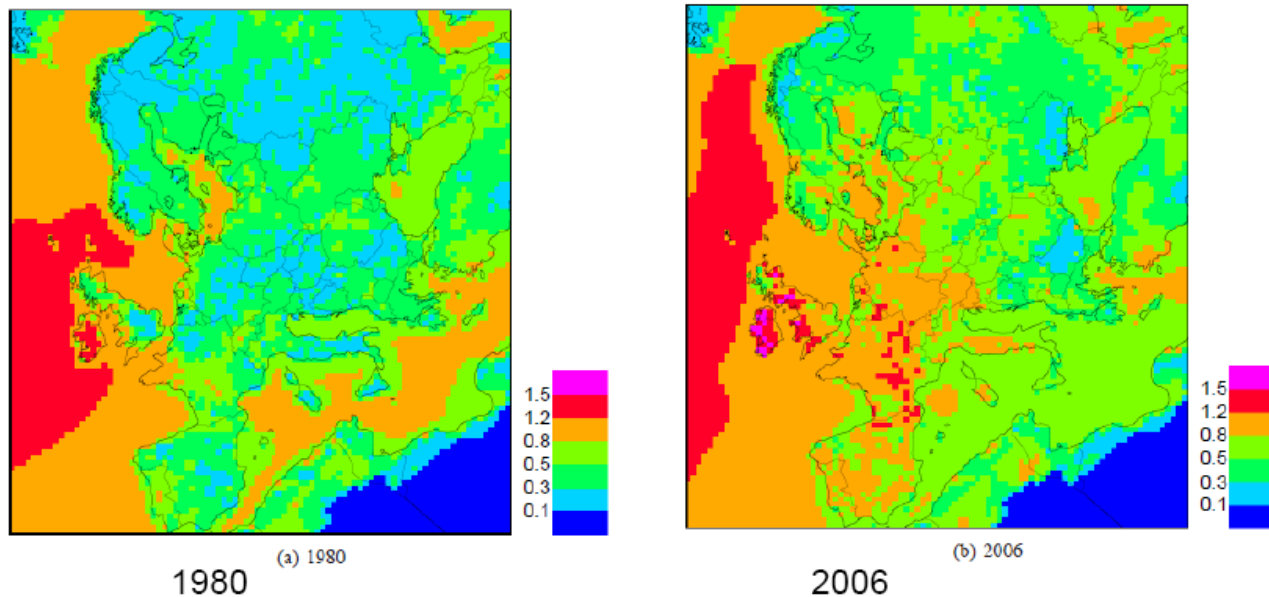
# Long-term trends in dry deposition

## % of UK emissions deposited in UK



# Sulphur deposition and feedbacks

## Co-deposition of SO<sub>2</sub> and NH<sub>3</sub>



1980

2006

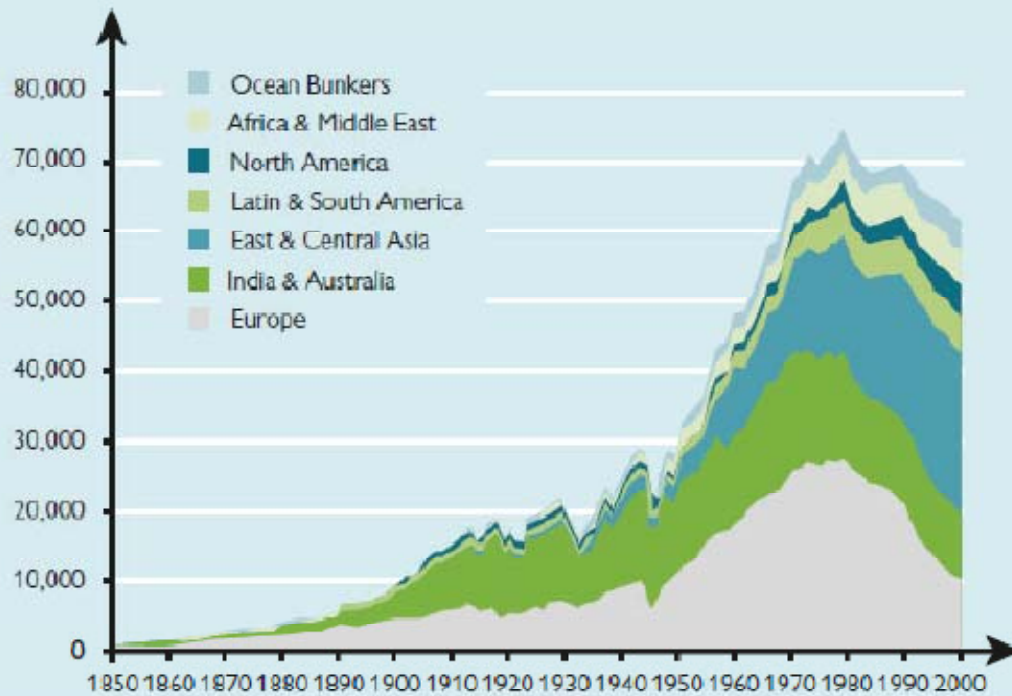
Modelled dry deposition velocities ( $\text{cm s}^{-1}$ ) of SO<sub>2</sub> over Europe for 1980 and 2006

**Increasing dry deposition velocity of SO<sub>2</sub> – less SO<sub>2</sub> available for sulphate formation – less negative RF from aerosols – positive feedback**

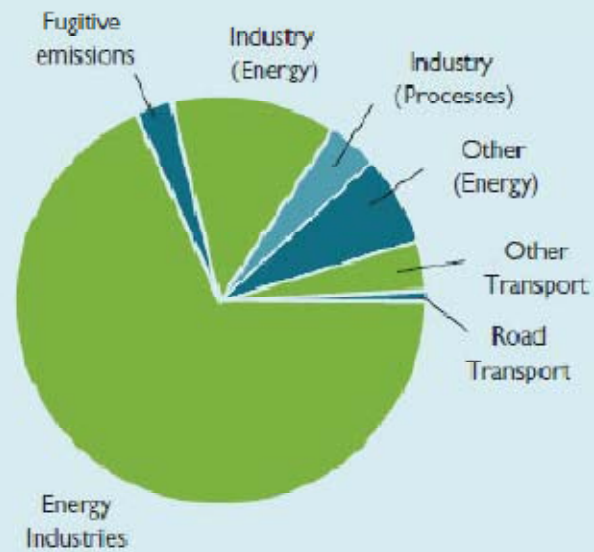
(Hilde Fagerli et al. in preparation, EMEP MSCW, 2009)

# Global SO<sub>2</sub> emissions by region, 1850-2000

SO<sub>2</sub> emissions  
1000 tonnes S



a)



SO<sub>2</sub> emissions by sector

b)

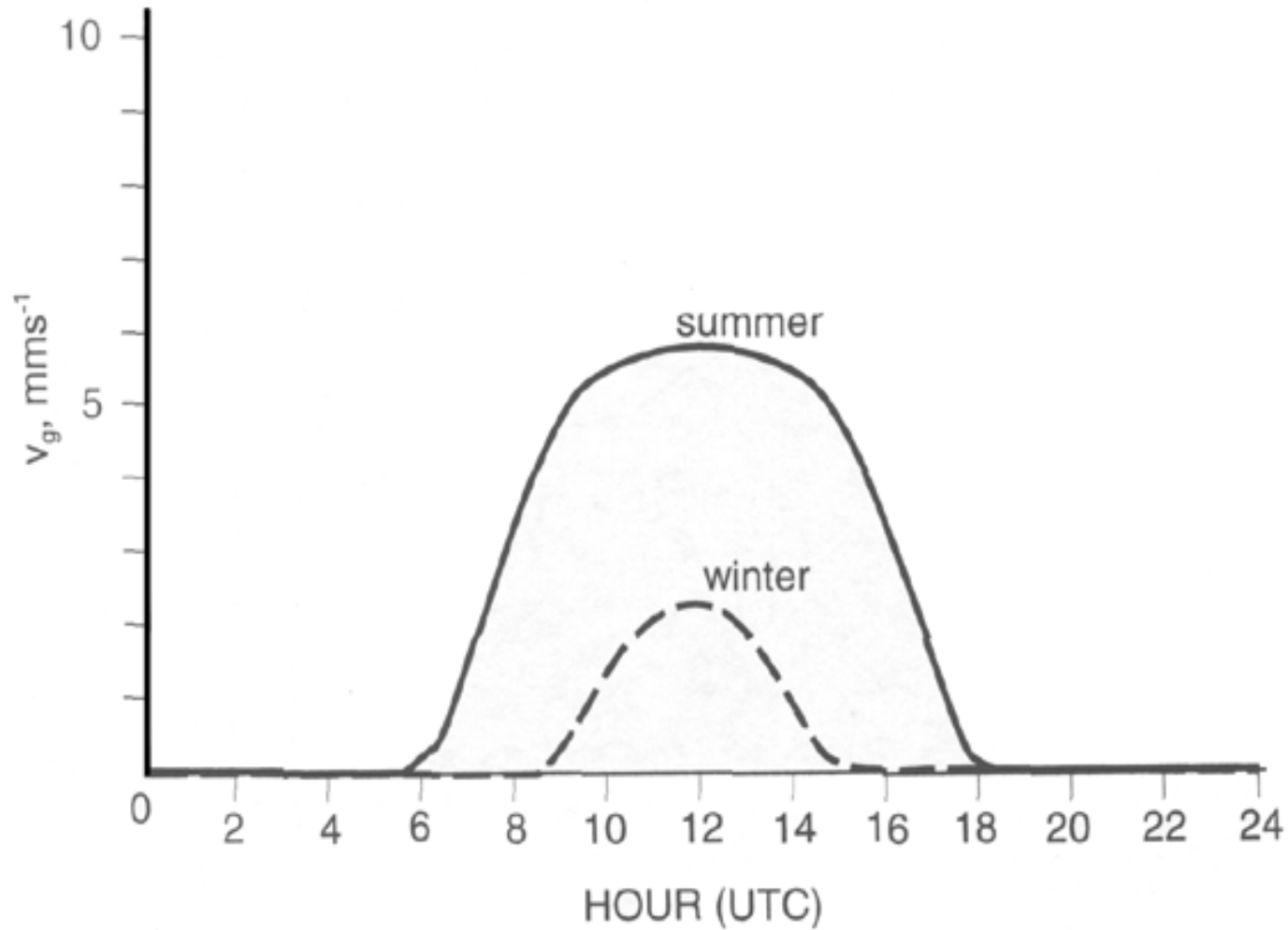
Source: a) Smith et al., 2004; b) EU-27 EEA

# Reactive nitrogen compounds

- **Oxidized Nitrogen NO, NO<sub>2</sub>, HNO<sub>3</sub>, HONO, PAN**
- Processes in soils
- Compensation points
- Canopy- atmosphere interactions and a Discussion of the canopy uptake of the fraction of soil NO emission converted to NO<sub>2</sub> and taken up by the canopy.
- Up-scaling of NO emission to regions, continents and global scales and their uncertainties.
- HNO<sub>3</sub> deposition, the presence of a surface resistance
- Uptake of NO<sub>2</sub> by vegetation and soil.
- Formation of HONO at terrestrial surfaces

# NO<sub>2</sub> deposition on vegetation

Diurnal cycle in NO<sub>2</sub> deposition velocity.



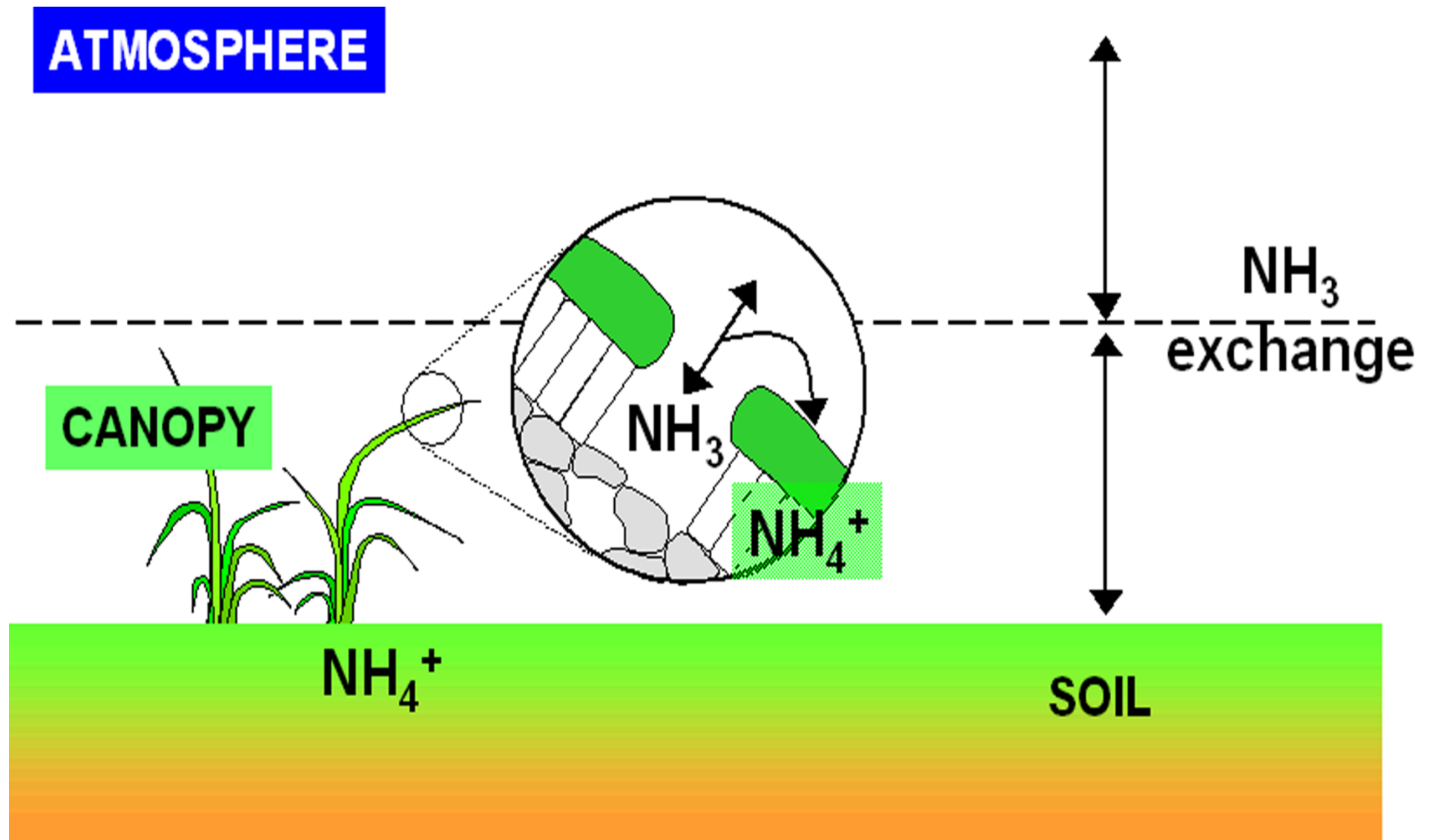
Ammonia

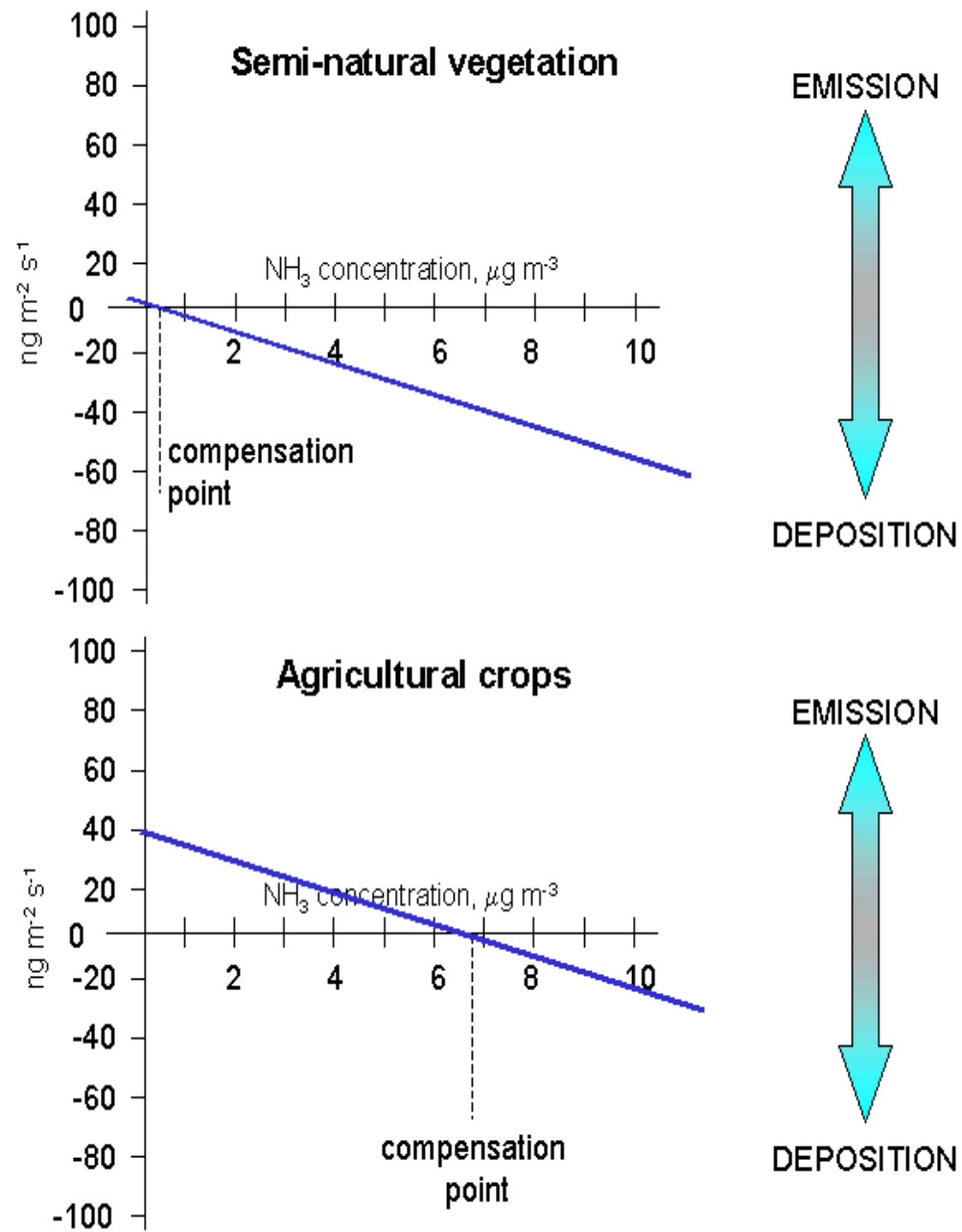
## Reduced Nitrogen ( $\text{NH}_3$ )

- Processes at the surface regulating  $\text{NH}_4^+$  concentrations in the apoplast, and their interaction with the atmosphere and external (leaf surface) factors.
- Compensation points
- Concentration dependent canopy resistances for  $\text{NH}_3$  deposition
- Up-scaling to the canopy, landscape, region and continent.

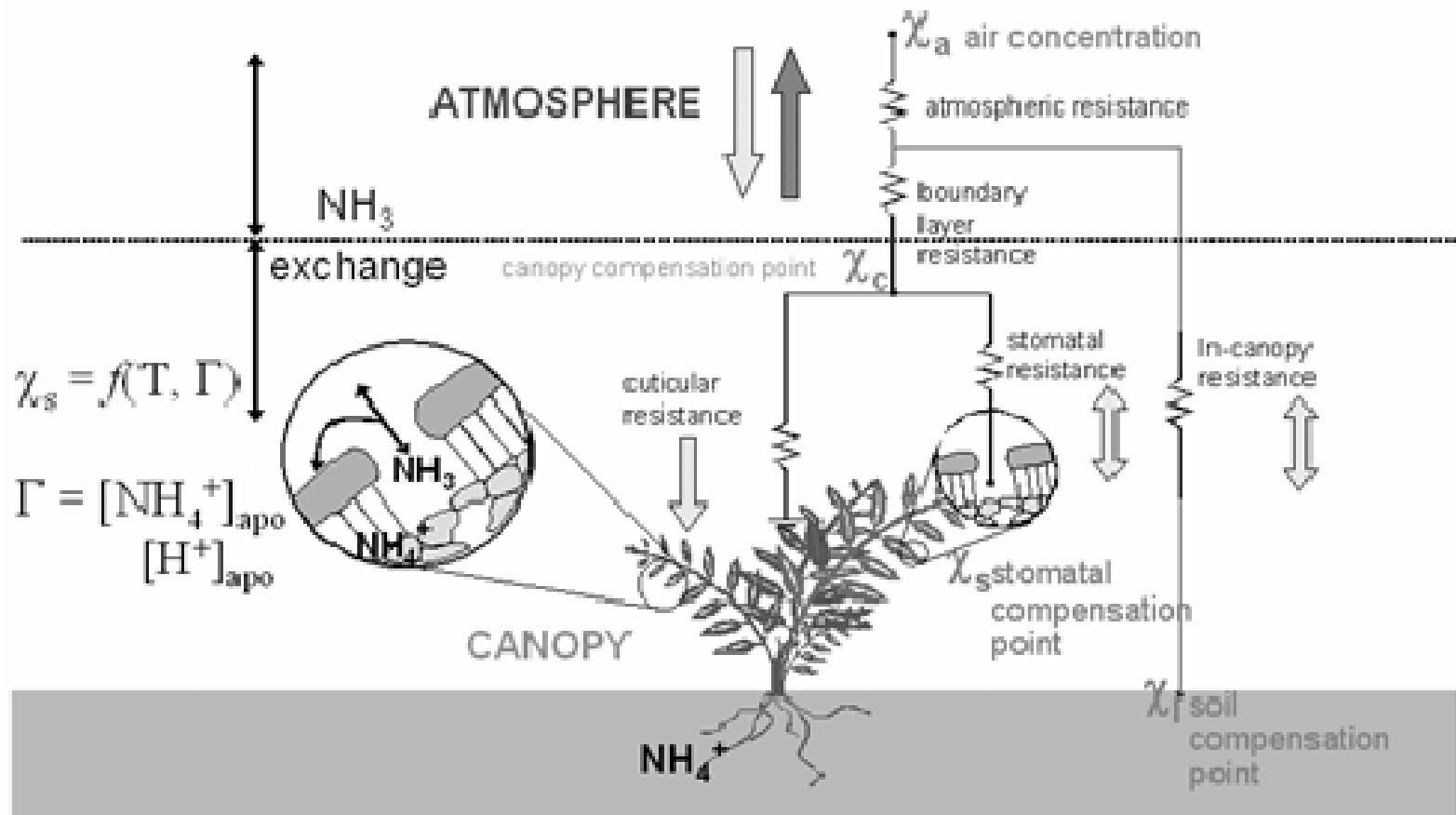


# Bi-directional exchange



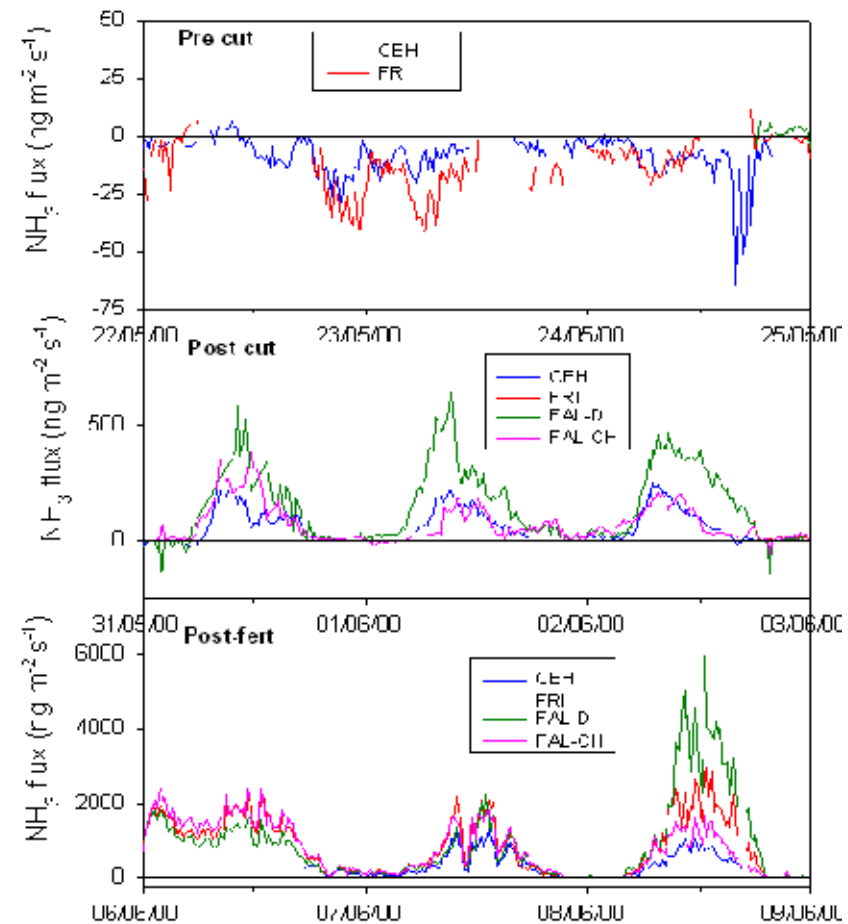


# Canopy compensation point model

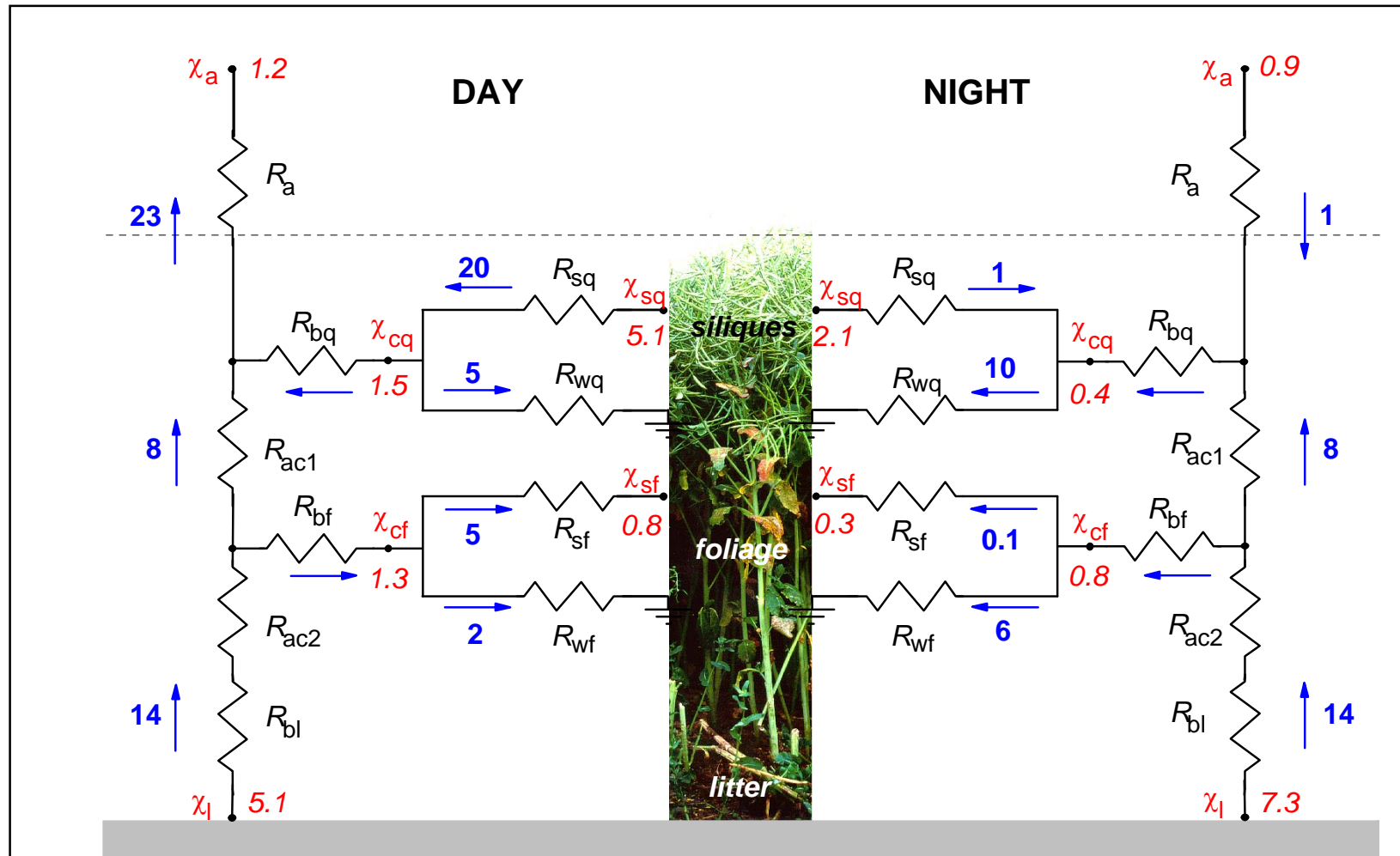


# Ammonia exchange over vegetation

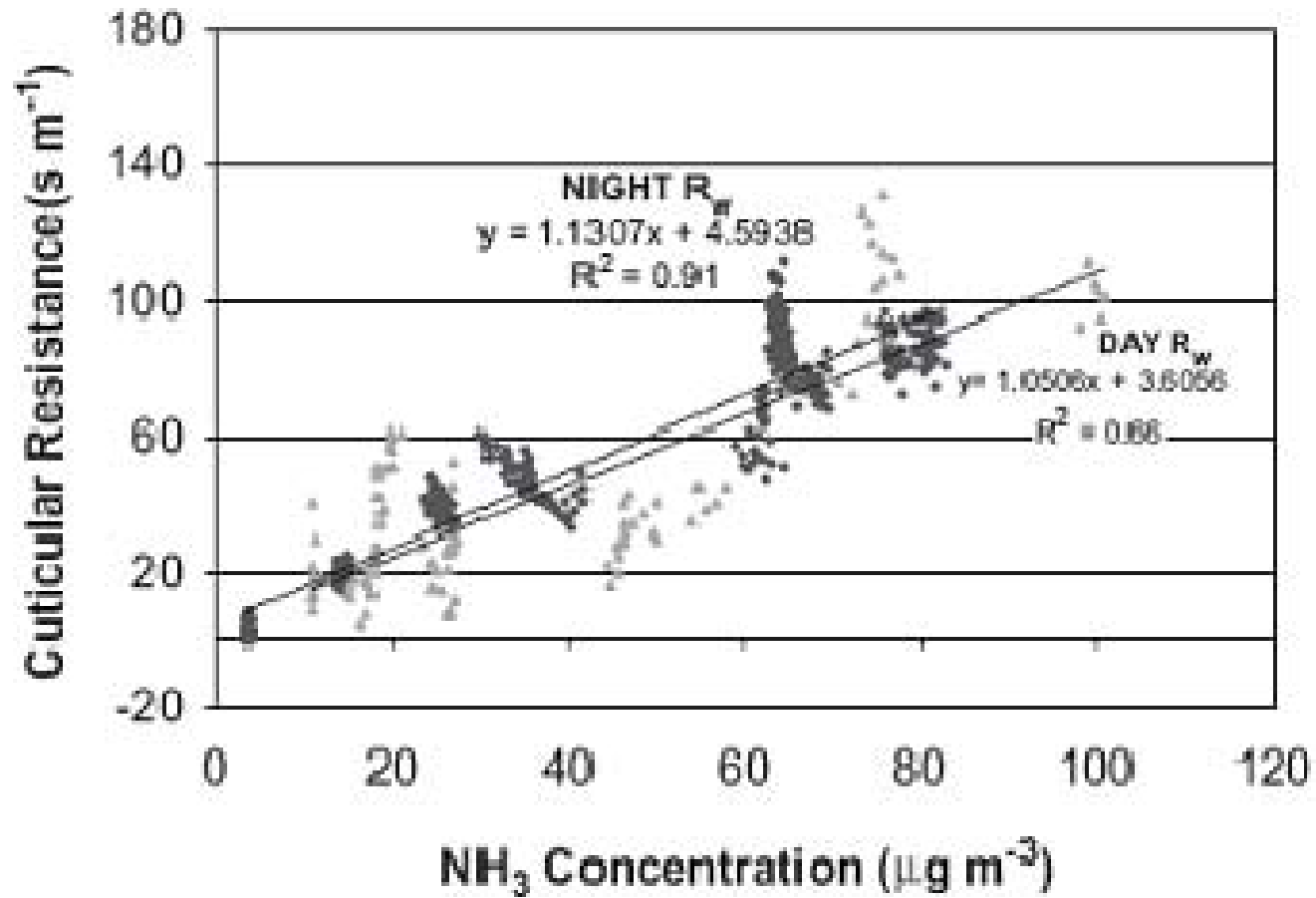
Intercomparison of methods to measure canopy-atmosphere exchange fluxes of  $\text{NH}_3$  at Braunschweig over agricultural grassland



# Resistance modelling of component fluxes of ammonia within plant canopies



## Concentration dependence of leaf surface uptake of ammonia

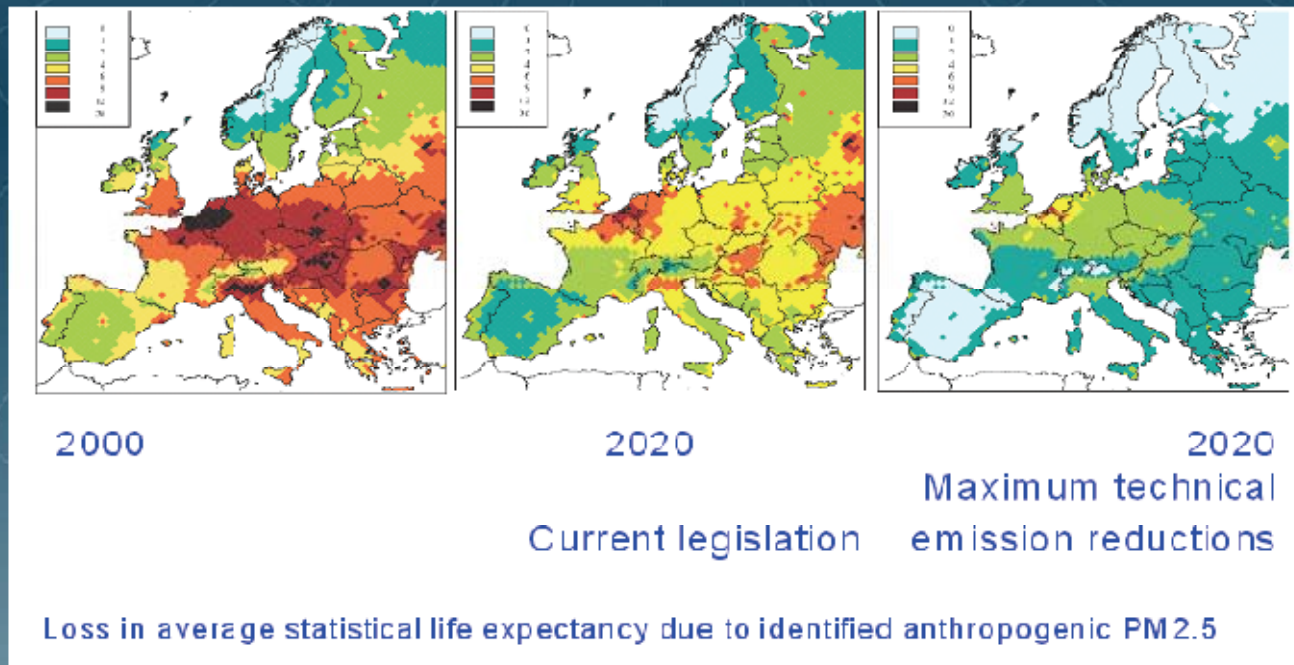


Aerosols

Particulate matter

# Particle emission and deposition

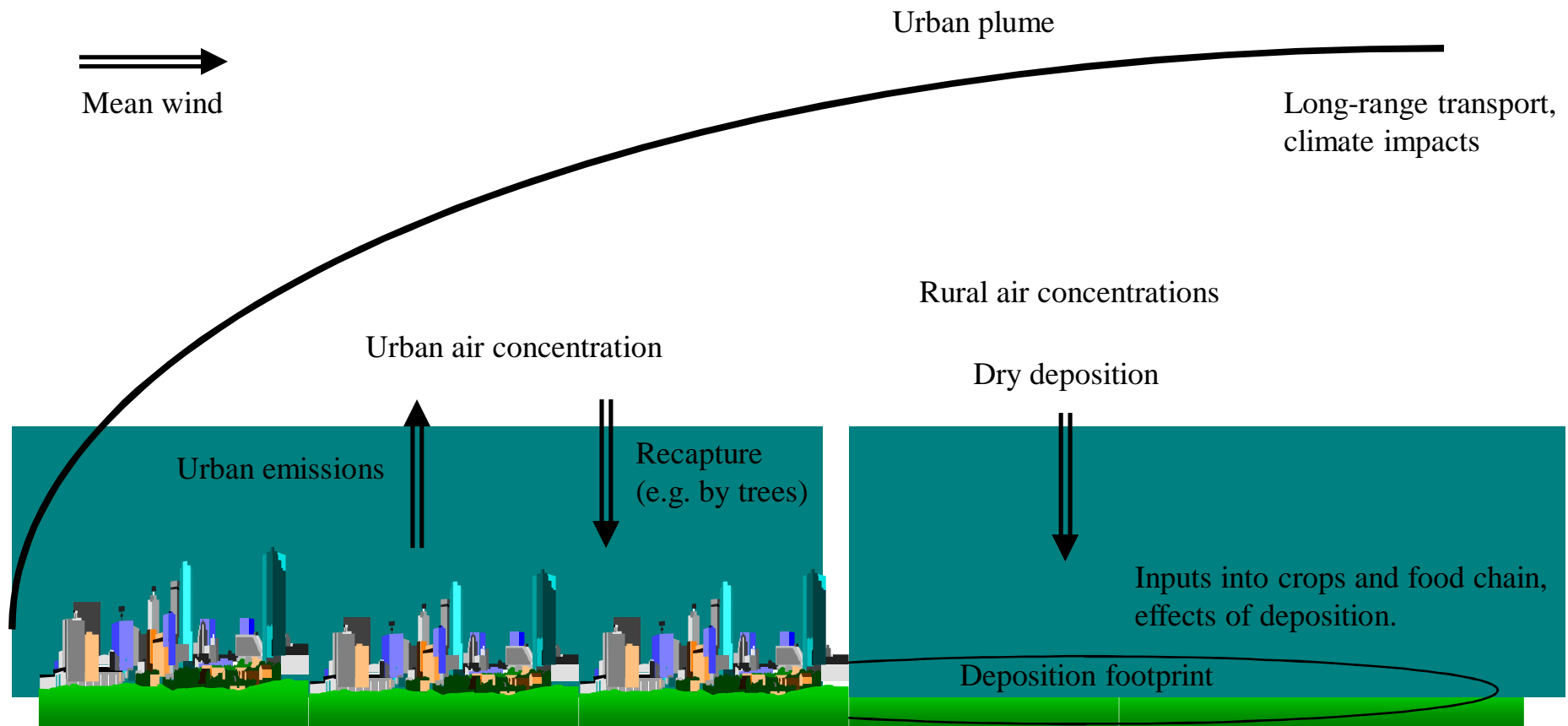
## Loss in life expectancy attributable to fine particles (months)



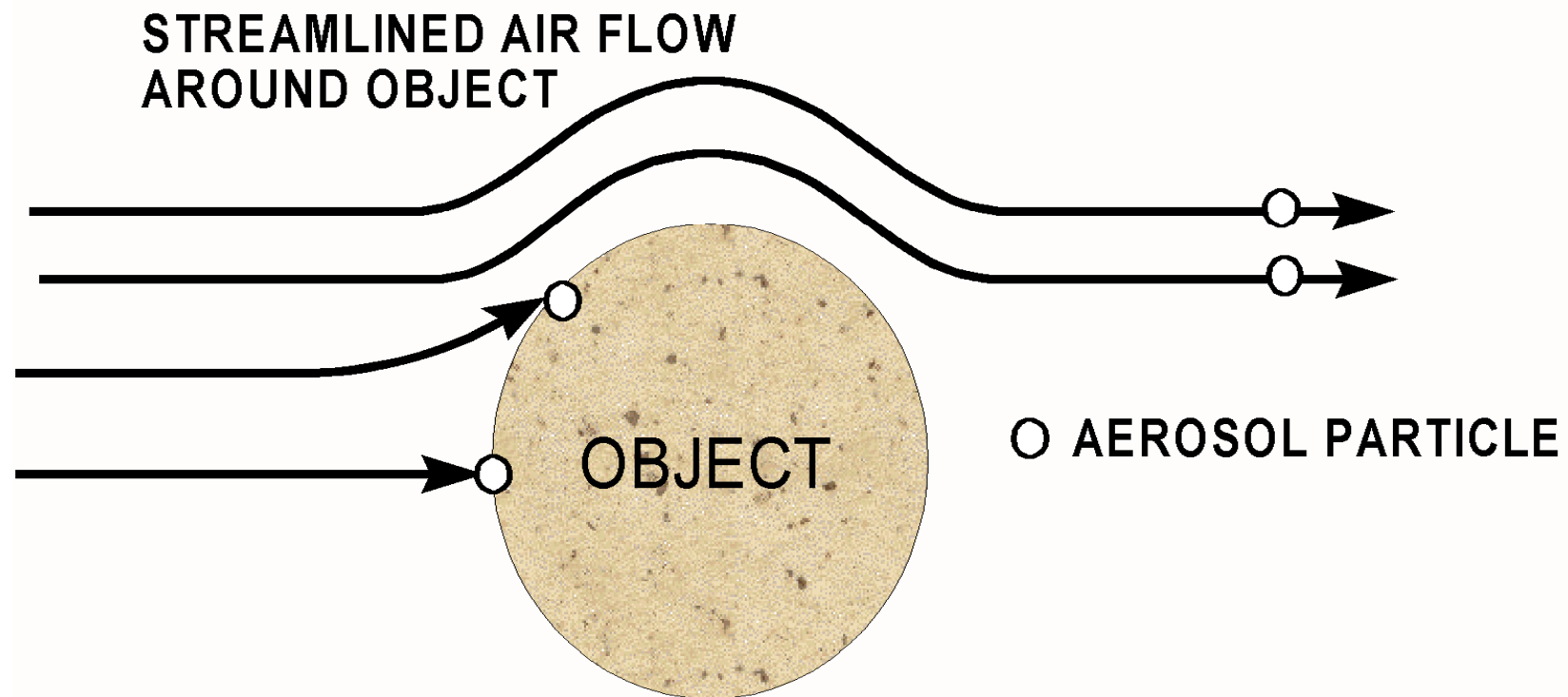
Source: IIASA



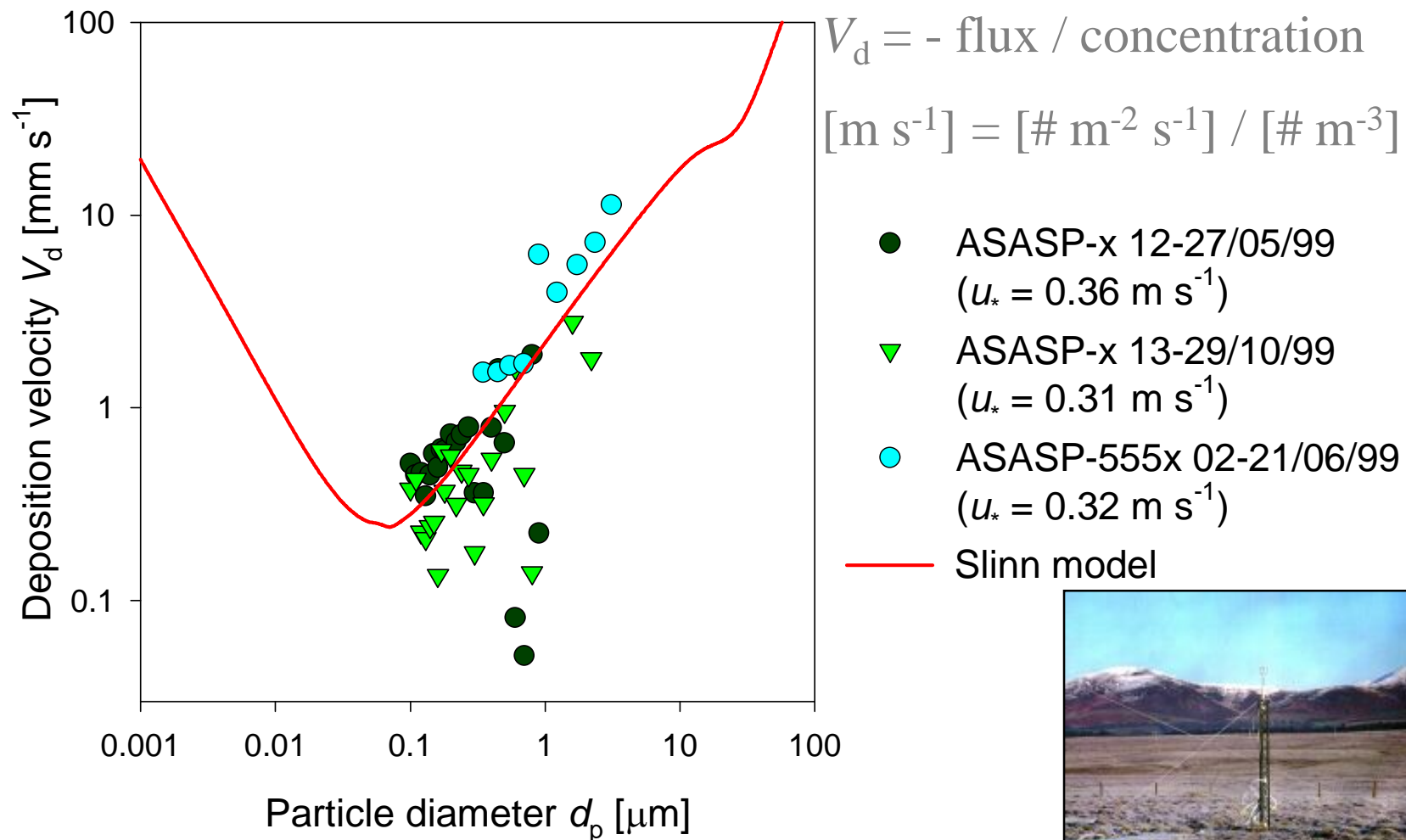
# Emission, recapture and deposition



# AEROSOL IMPACTION

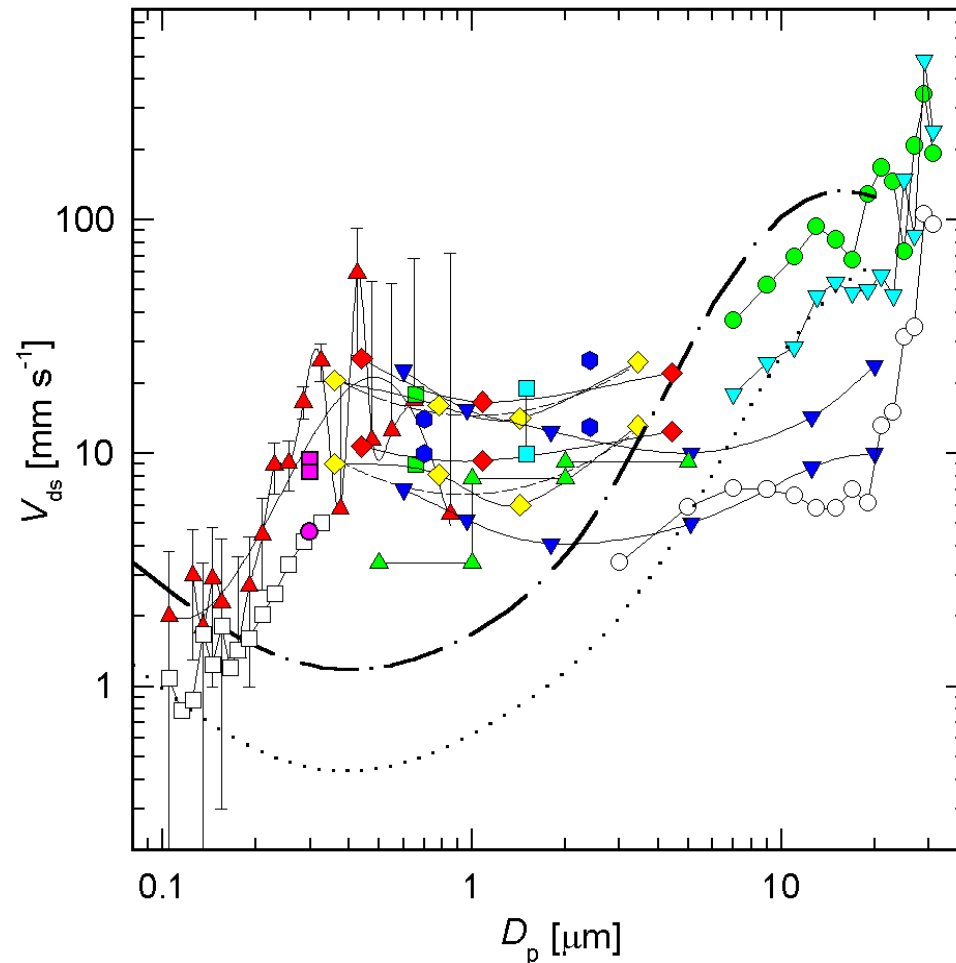


# Aerosol deposition velocities as a function of size to moorland vegetation



# Aerosol deposition velocities as a function of size to trees

- ◆ Höfken and Gravenhorst (1983)
- ◆ Grosch and Schmitt (1988)
- Brückman (1988)
- Joutsenoja (1992)  $u^* = 0.68 \text{ m s}^{-1}$ ,  
 $u = 5.6 \text{ m s}^{-1}$ ,  $z_0 = 0.08 - 0.13 \text{ m}$
- ▼ Waraghai and Gravenhorst (1989)
- ▼ Gallagher *et al.* (1992)  
 $u^* = 0.71 \text{ m s}^{-1}$ ,  $z_0 = 15.7 \text{ cm}$
- Gallagher *et al.* (1992)  
 $u^* = 0.75 \text{ m s}^{-1}$ ,  $z_0 = 18.6 \text{ cm}$
- Beswick *et al.* (1991)  
 $u^* = 0.37 \text{ m s}^{-1}$ ,  $z_0 = 30.1 \text{ cm}$
- ▲ Gallagher *et al.* (1997a)
- ▲ Lorenz and Murphy (1989)  
Gradient Loblolly Pine,  $h_c = 9 \text{ m}$
- Höfken *et al.* (1983) TF Beech
- Höfken *et al.* (1983) TF Spruce
- ⋯ Slinn (1982),  $u^* = 0.65 \text{ m s}^{-1}$   
Unmodified Model Result
- Slinn (1982),  $u^* = 1.30 \text{ m s}^{-1}$   
Unmodified Model Result
- Choubedar and Fowler (2000)  
 $^{210}\text{Pb}$  grassland Rothamsted ( $u_* = 0.14 \text{ m s}^{-1}$ )
- Choubedar and Fowler (2000)  
 $^{210}\text{Pb}$  forests Rothamsted ( $u_* = 0.36 \text{ m s}^{-1}$ )





## Measurements at the Nelson Monument



# Example footprints and wind sector dependence of the flux

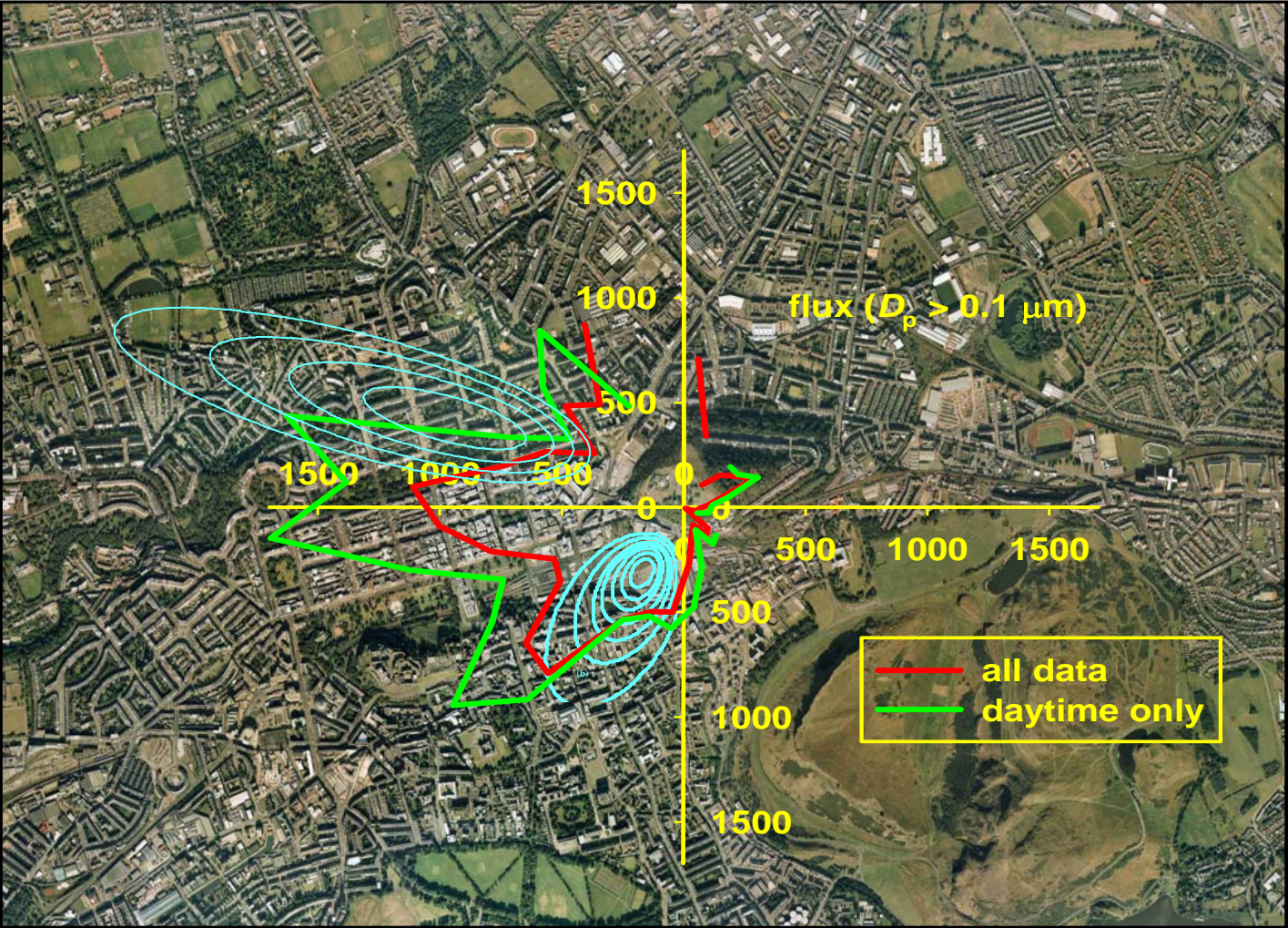


Image copyright Cities Revealed®, ©The GeoInformation Group, 1998

Probably stop here!



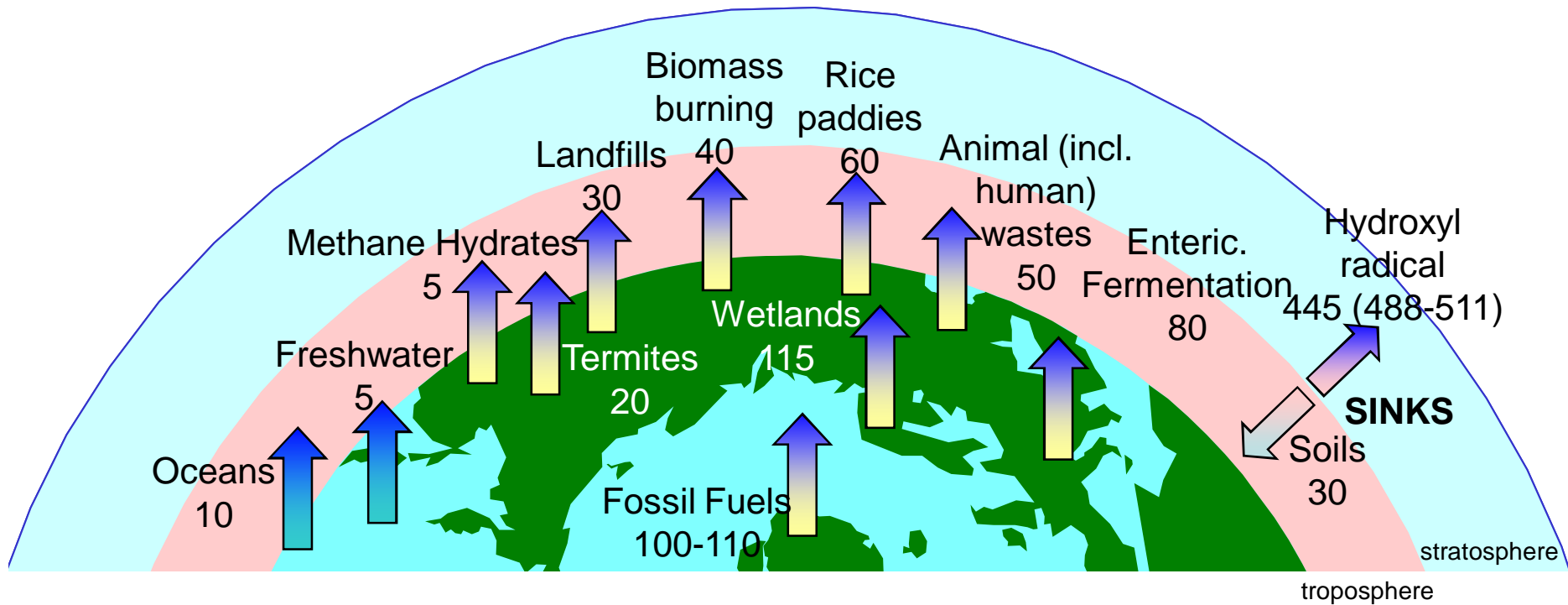


# Methane

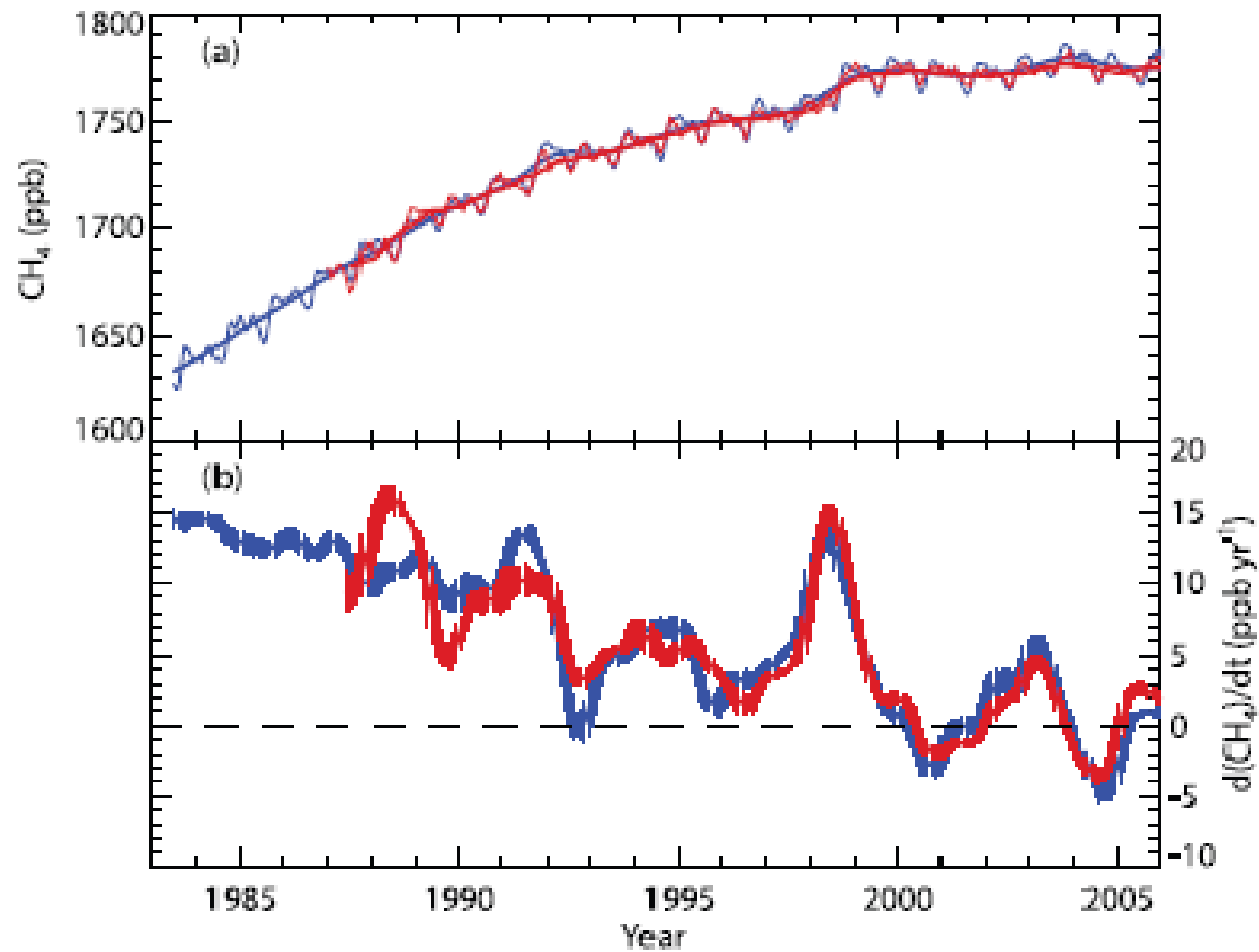
- Bi-directional,
- methane uptake by bacteria in aerobic soils
- Methane emission (by methanogenic bacteria) in anaerobic conditions.

## GLOBAL METHANE BUDGET (Mt)

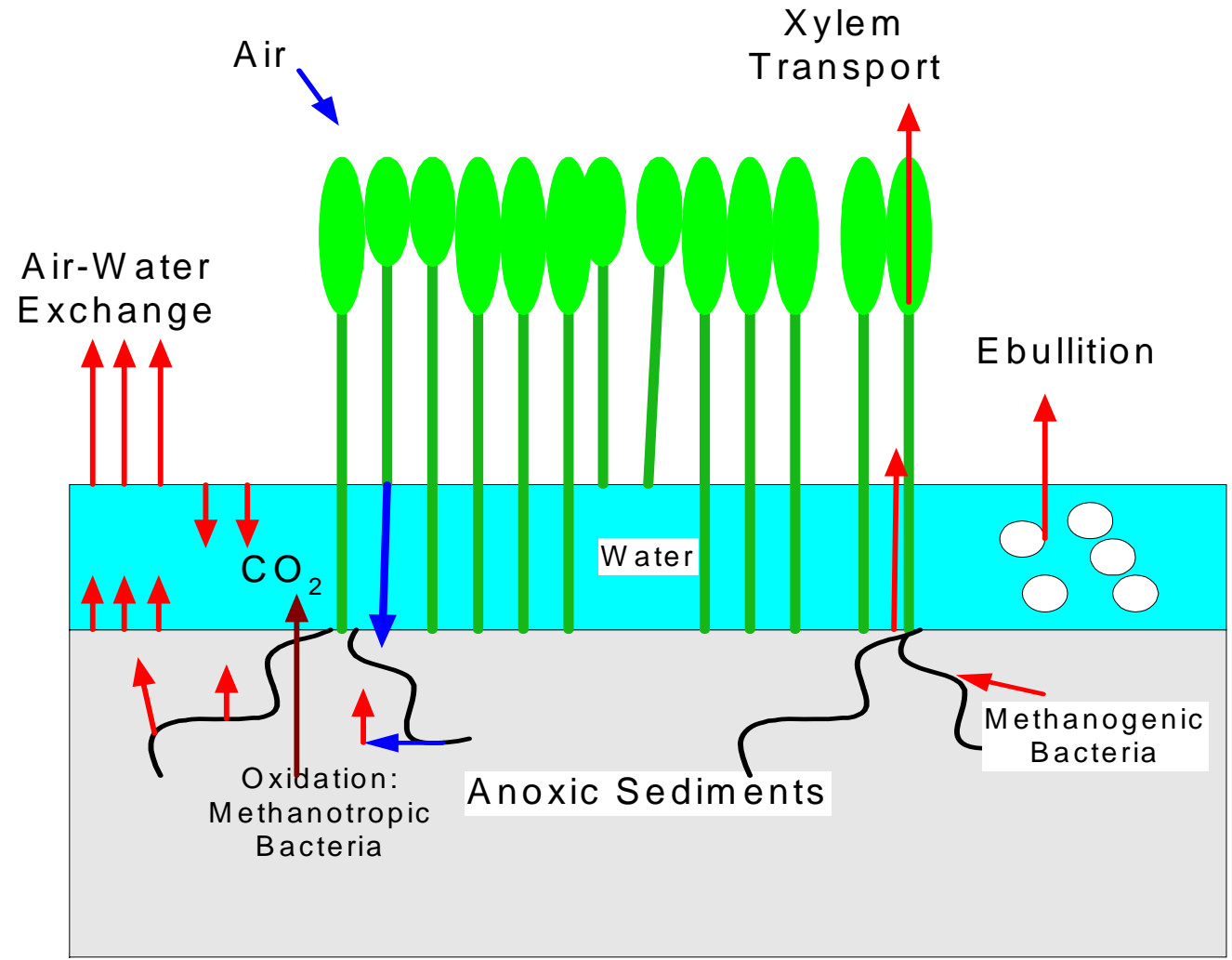
Annual Source	515 Mt
Annual Sink	475 Mt
Net Input to Troposphere	30 - 50 Mt



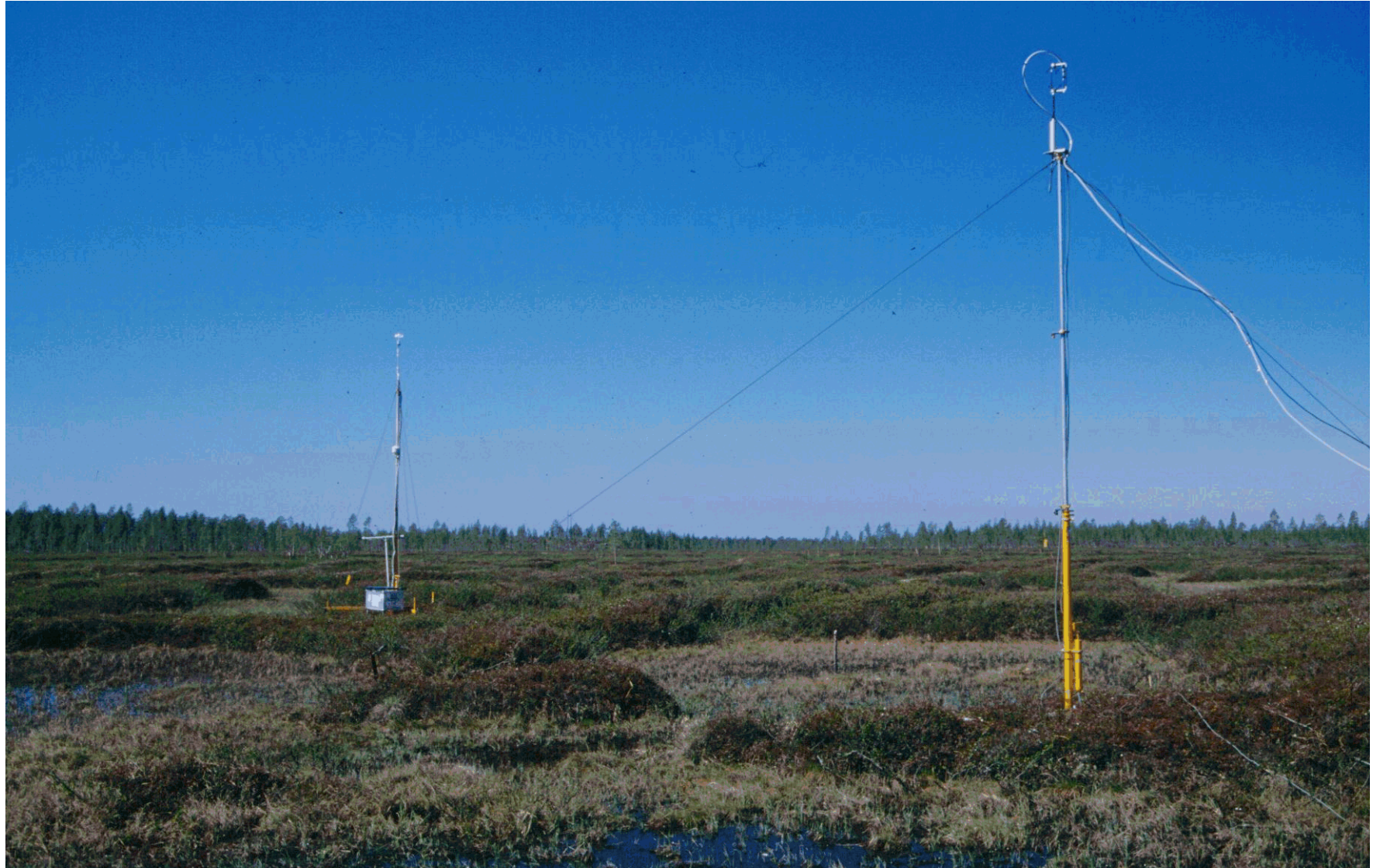
# Greenhouse gas exchange and ecosystems: methane



# Fluxes, Sources and Sinks of Methane

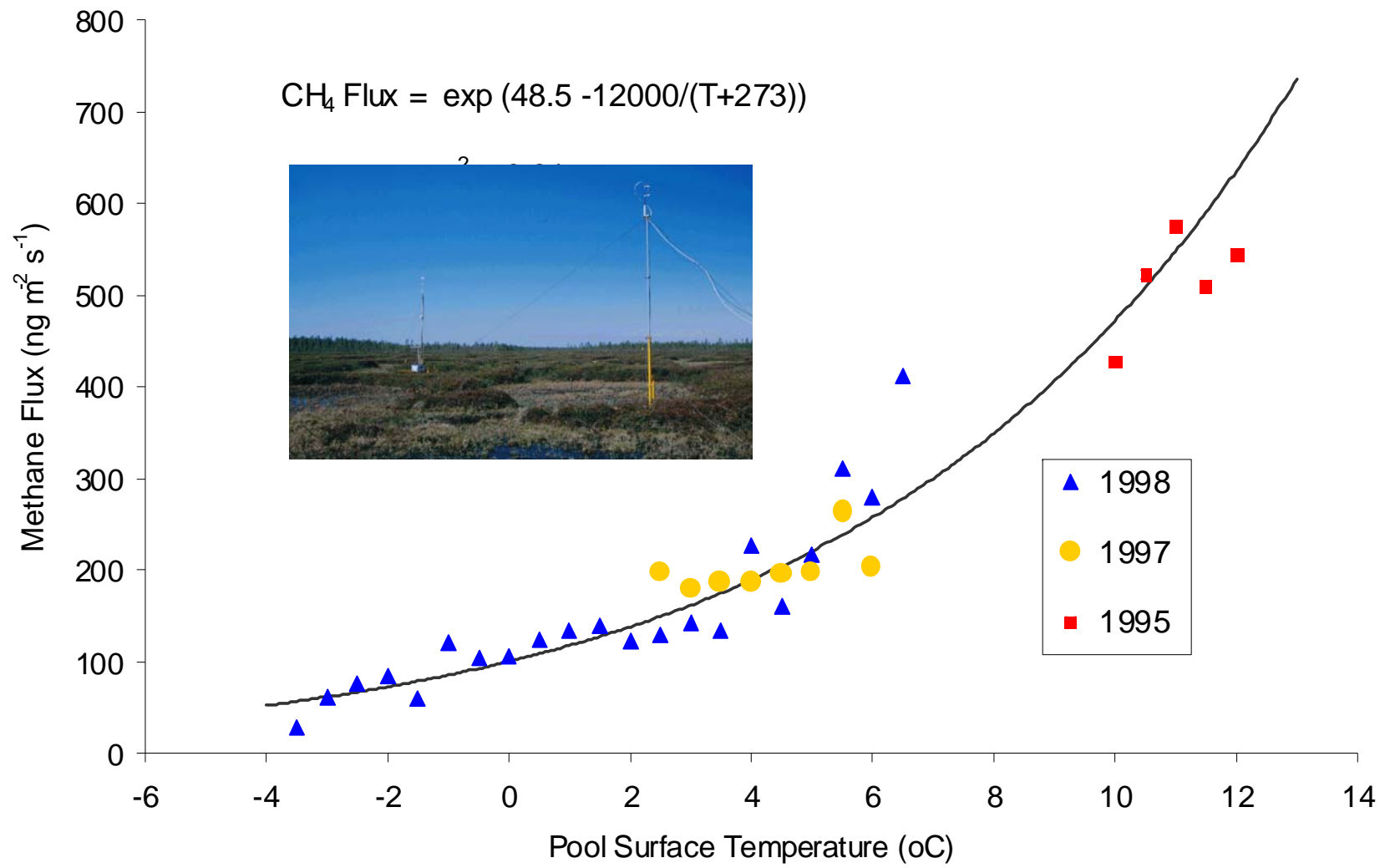


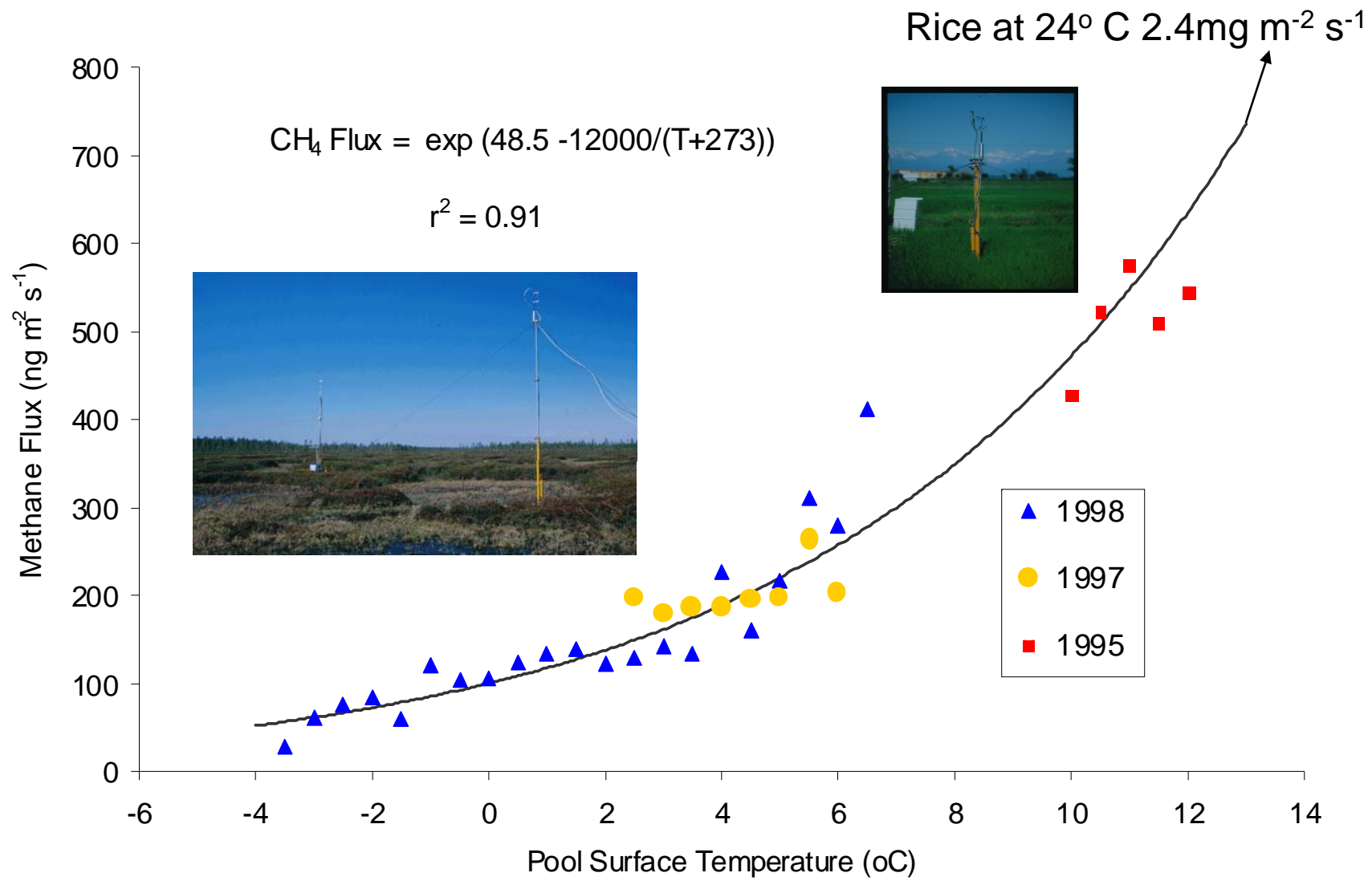
# Lapland methane flux measurements





Source: CEH

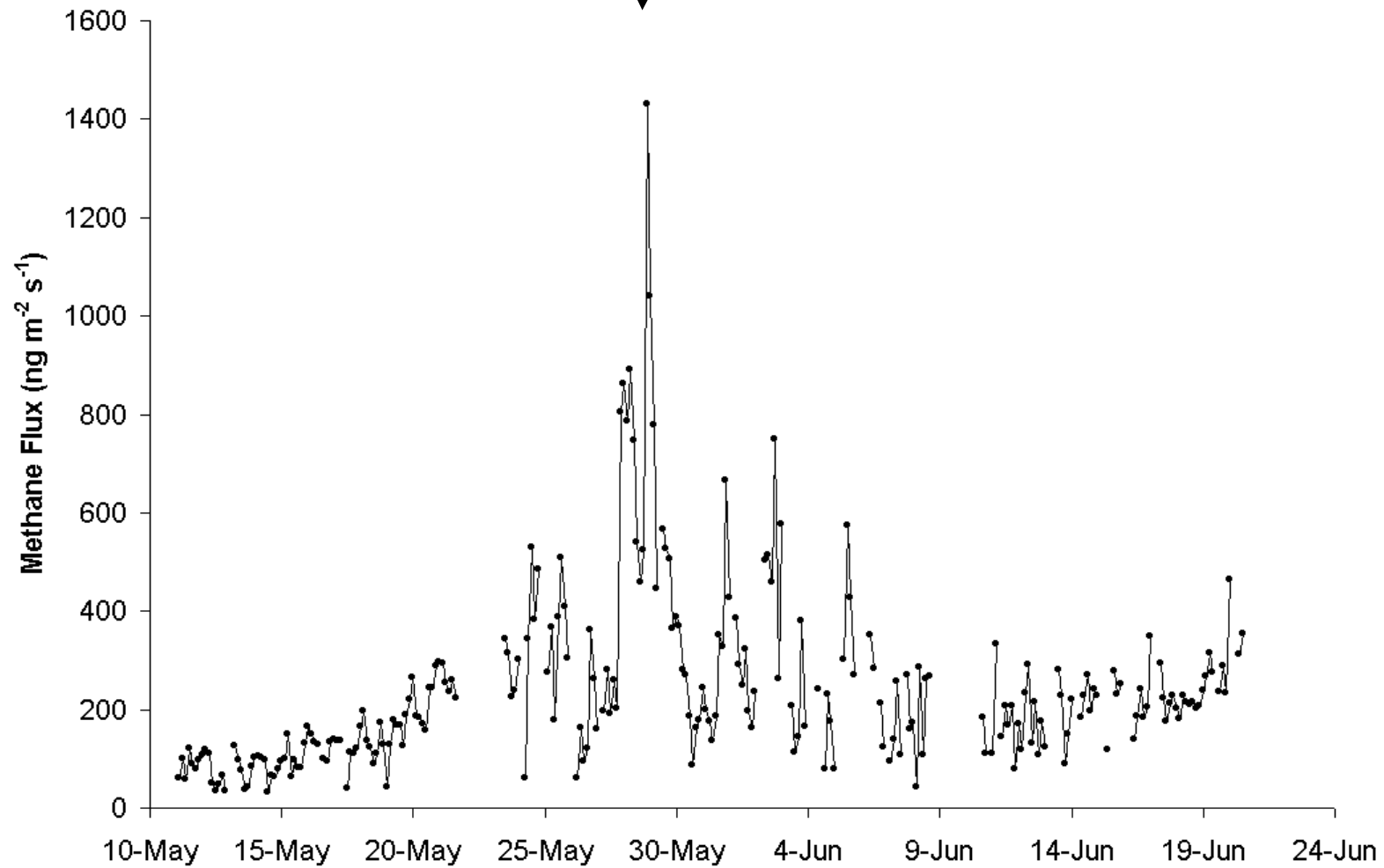




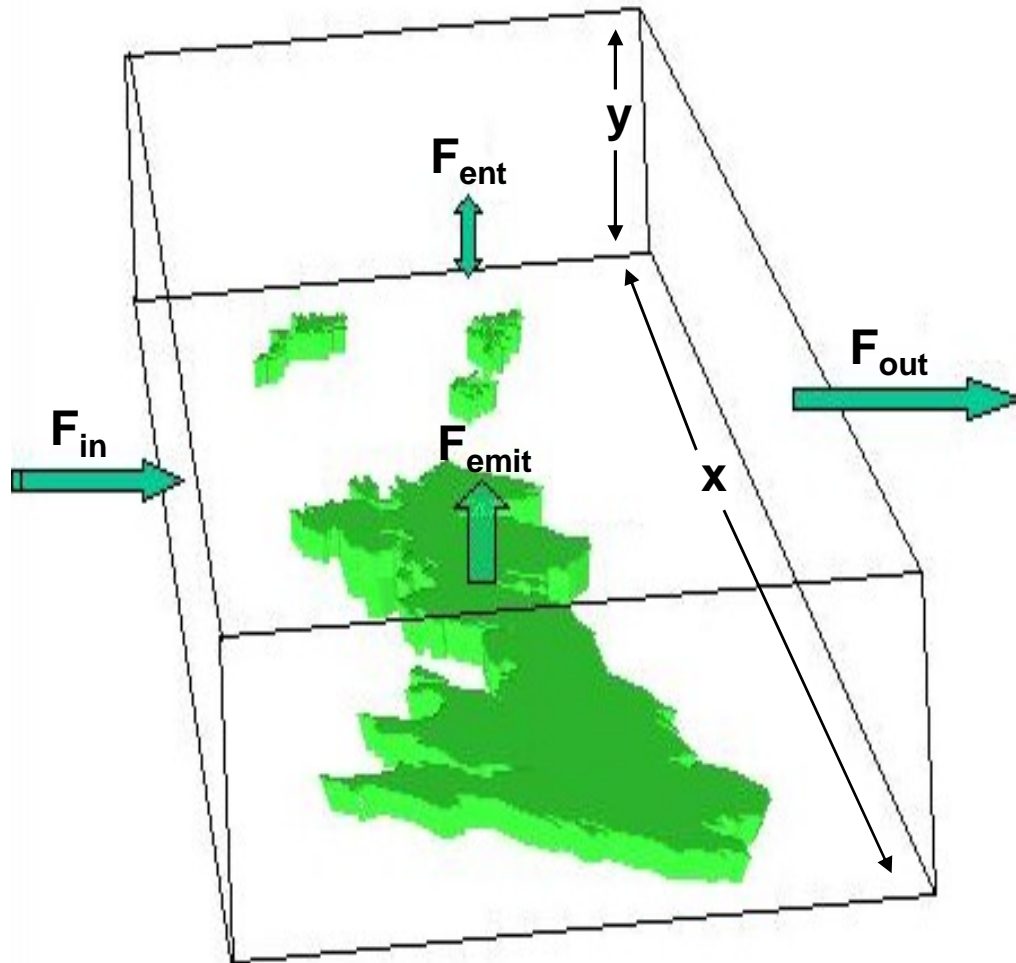


# Methane emission during the snow-melt period in Lapland

Snow-melt



# Estimating Annual Emissions – Box Technique



- $F_{\text{emit}} = F_{\text{out}} - F_{\text{in}} - F_{\text{ent}}$
- $F_{\text{in}} = \iint u_{\text{in}}(y,z)c(y,z)dydz$
- $F_{\text{out}} = \iint u_{\text{out}}(y,z)c(y,z)dydz$
- $F_{\text{ent}} = 0$

$F_{\text{emit}}$  = Flux emitted

$F_{\text{in}}$  = Flux in

$F_{\text{out}}$  = Flux out

$u$  = wind speed

$c$  = concentration

# Atmospheric Mass Budgets at Regional Scales

## UPWIND

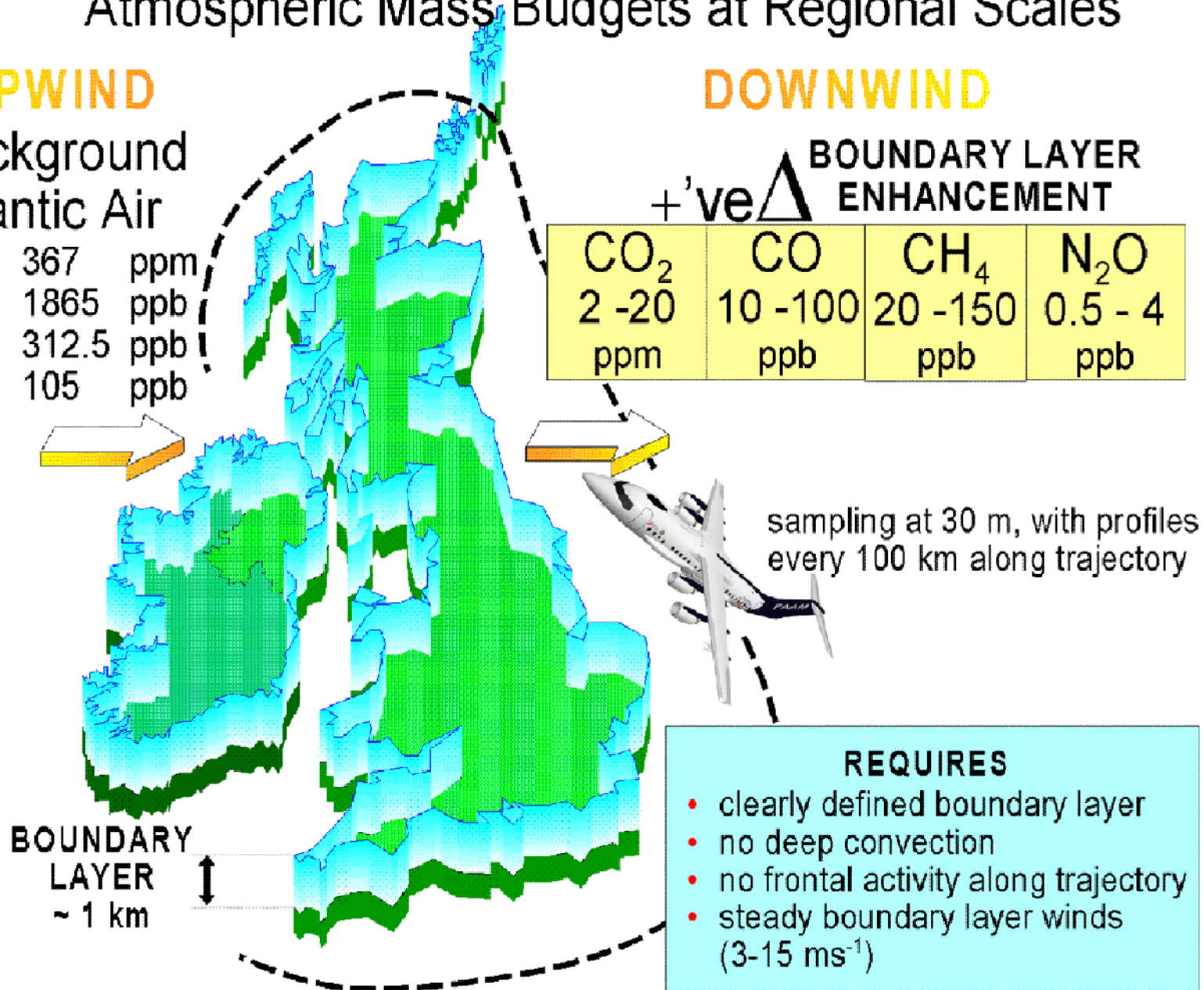
Background  
Atlantic Air

CO <sub>2</sub>	367	ppm
CH <sub>4</sub>	1865	ppb
N <sub>2</sub> O	312.5	ppb
CO	105	ppb

## DOWNWIND

+ve  $\Delta$  BOUNDARY LAYER  
ENHANCEMENT

CO <sub>2</sub>	CO	CH <sub>4</sub>	N <sub>2</sub> O
2 -20	10 -100	20 -150	0.5 - 4
ppm	ppb	ppb	ppb



BOUNDARY  
LAYER  
~ 1 km

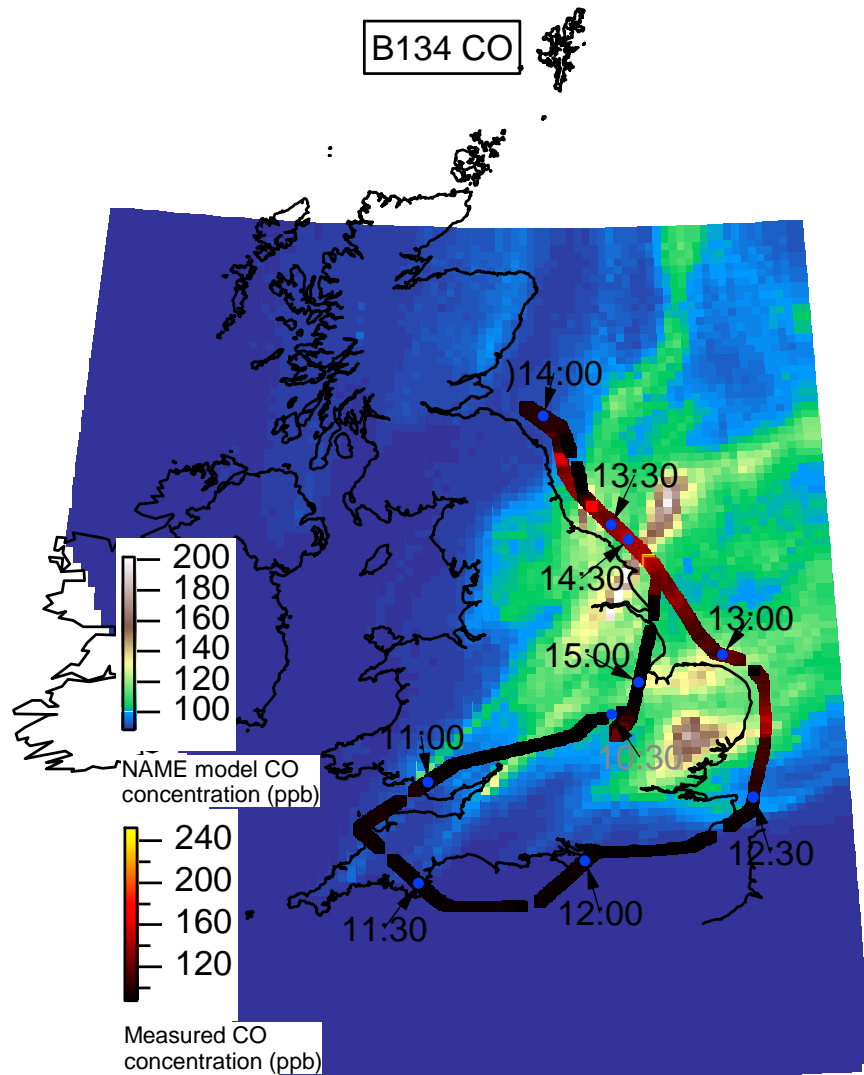
- REQUIRES**
- clearly defined boundary layer
  - no deep convection
  - no frontal activity along trajectory
  - steady boundary layer winds (3-15 ms<sup>-1</sup>)

# UK Facility Airborne Atmospheric Measurement (FAAM): BAe-146

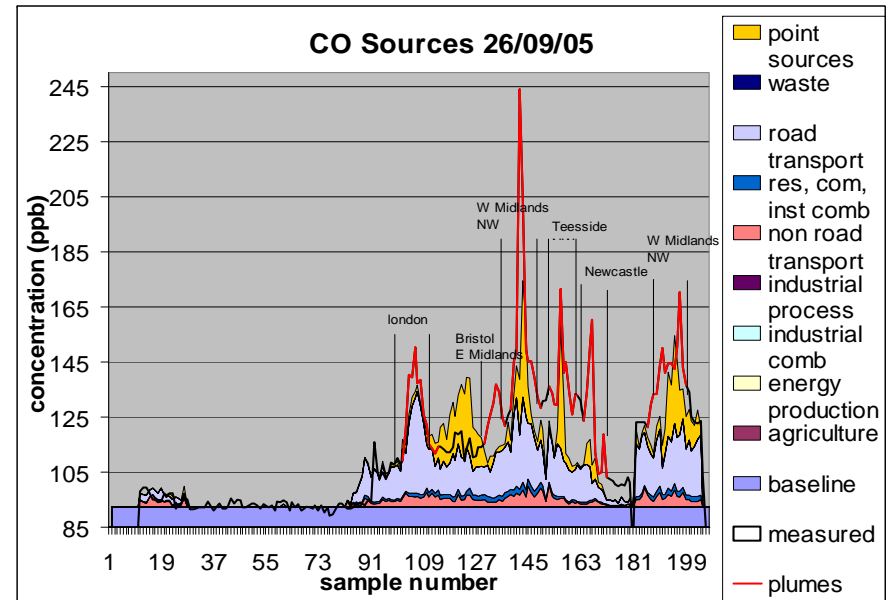
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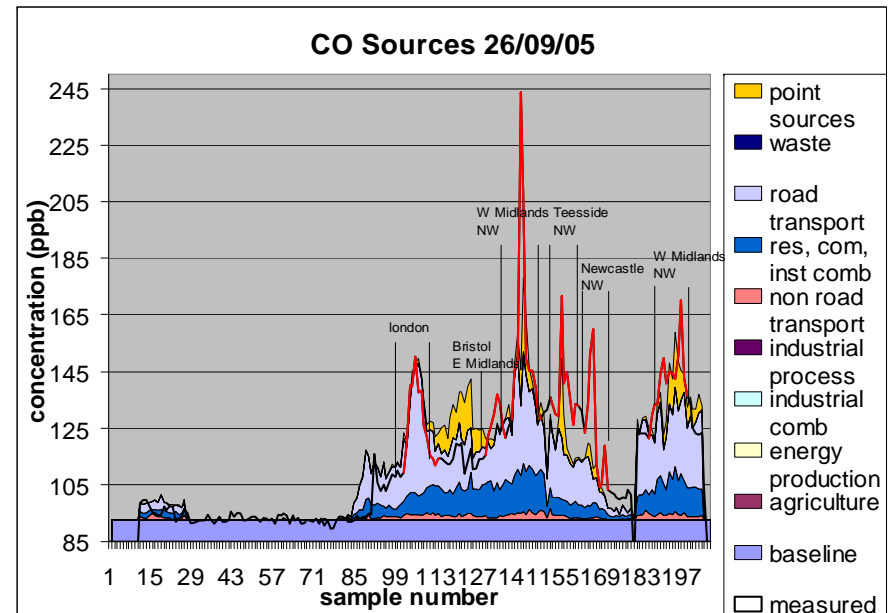
**Flight duration:** max. 5:30 hours



### Original emissions



### Adjusted emissions



## Results: NAME Sectorial Emission Adjustment Technique

	CO (kT/y)	CO <sub>2</sub> (kT/y)	N <sub>2</sub> O (kT/y)	CH <sub>4</sub> (kT/y)
<b>NAEI (2003)</b>	<b>2757</b>	<b>572196</b>	<b>130</b>	<b>1933</b>
B92	1900 ± 147	490000 ± 68500	880 ± 456	*8700 ± <sup>7770</sup> / <sub>6870</sub>
B97	3700 ± 605	470000 ± 117000	540 ± 427	4600 ± 1190
B102	2700 ± 250	410000 ± 93000	620 ± 334	*****
B111	*3000 ± 202	*1200000 ± 181000	*930 ± 223	*****
B112	1700 ± 292	620000 ± <sup>222000</sup> / <sub>130000</sub>	*2200 ± <sup>1260</sup> / <sub>239</sub>	9600 ± 3440
B113	2800 ± 847	840000 ± 319888	*2900 ± 893	12000 ± 6910
B118	800 ± 141	650000 ± 87266	320 ± 199	*****
B119	1900 ± 102	510000 ± 145000	190 ± 106	*****
B126	3300 ± 66	620000 ± 130000	3800 ± <sup>3760</sup> / <sub>3938</sub>	9300 ± 2730
B130*	2400 ± 48	420000 ± 82800	200 ± 72	1900 ± <sup>661</sup> / <sub>231</sub>
B132*	1600 ± 43	330000 ± 127000	390 ± 64	1600 ± 210
B134	4700 ± 167	510000 ± 75200	320 ± 186	6500 ± <sup>4070</sup> / <sub>3510</sub>
B136	2900 ± 334	580000 ± 246000	620 ± 532	*2800 ± <sup>1570</sup> / <sub>1390</sub>
<b>AVERAGE</b>	<b>2400 ± 226</b>	<b>560000 ± <sup>139000</sup>/<sub>133000</sub></b>	<b>350 ± 208</b>	<b>4000 ± <sup>1400</sup>/<sub>1290</sub></b>