

2465-2

**Joint ICTP-IAEA Workshop on Evaluating Groundwater Pathways and  
Residence Times as part of Site Investigations and Post-Closure Safety  
Assessments for Geological Repositories'**

*17 - 21 June 2013*

**Shallow Active Systems Analysis and Recent Advances in Isotopic Methods for  
Residence Time Determinations <50,000y. "DATING GROUNDWATER OF  
AGES YOUNGER THAN 50,000 YEARS**

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## Dating groundwater of ages younger than 50,000 years

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## Young ground water dating

- Young groundwater is typically found at depths from 0 to 50 meter in unconsolidated sediments and at depths up to 300 meter in fractured-rock systems.
- Shallow groundwater systems are commonly used for drinking water sources and they make up a large part of the baseflow in rivers and lakes.
- Because they are shallow (= recently recharged), they are more susceptible to contamination than deeper groundwaters.
- Groundwater age can be used to determine recharge rates and refine hydrologic models of groundwater systems and thus to predict the contamination potential and estimate the time needed to flush contaminants through a groundwater system.
- The 0- to 50-year time scale is particularly relevant to environmentally sensitive shallow groundwater systems, but before 1980s, there were no reliable means of dating groundwater recharged during this time scale.

## Types of dating tools

1. Tracer Dating Technique: time vs. tracer's concentration curve to date groundwater age  
CFC & Tritium
2. Radioactive Decay Dating:  
 $^3\text{He}$ - $\text{T}$ ,  $^{14}\text{C}$ , Radio-Kryptons (half-life & concentration of daughter (+ parent) isotopes)
3. Linear Accumulation Dating: Using time-dependent input from external source  
 $^4\text{He}$  &  $^{40}\text{Ar}$  (ages from concentrations and known input rates).

## Tritium & CFCs

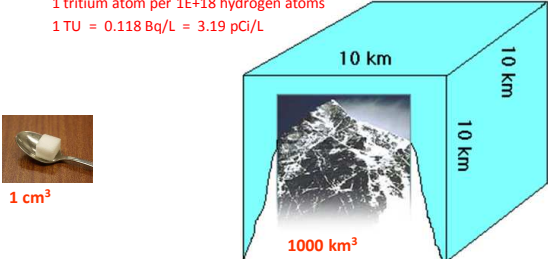
1. During the past 50 years, human activities have released an array of chemical substances to the atmosphere.
2. Tritium ( $^3\text{H}$ ) in water vapour from detonation of nuclear bombs in the 1950's and early 60'.
3. CFC (Chloro-Fluoro-Carbons) from refrigeration and other uses from 1950's to 80's, dissolve in precipitation.
4. These atmospheric substances became incorporated in the Earth's hydrologic cycle and can be found in groundwater that has been recharged within the past 50 years.

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## Tritium

- Tritium (Half-life 12 years)
  - Sources
    - Non-Human: Cosmic Ray bombardment on  $^{14}\text{N}$  in upper atmosphere (a fast neutron (which must have energy greater than 4.0 MeV) interacts with atmospheric nitrogen). Tritium concentrations due to this natural process is estimated to be between 4 to 25 TU depending on location

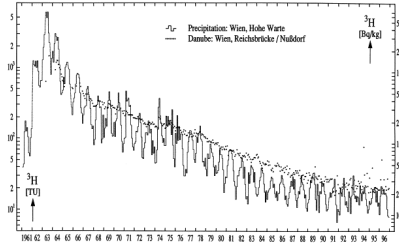
1 tritium atom per  $1\text{E}+18$  hydrogen atoms  
 1 TU =  $0.118 \text{ Bq/L} = 3.19 \text{ pCi/L}$



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## Tritium

- Tritium ( $^3\text{H}$ , Half-life 12 years)
  - Man-made Tritium: Released during 1950-60s nuclear bomb tests
  - >1000 TU in mid 60s in Northern hemisphere
  - Groundwater recharged after this time period is “labeled” with high Tritium signals.



The decline of tritium is not due only to decay (5.5 % per year), but also due to attenuation by the oceans and groundwaters has reduced levels considerably faster than by decay alone.

Theodorsson, Applied Radiation and Isotopes, 1999

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## Tritium

Tritium is a useful guide to distinguish between recharge that occurred before atomic weapons testing and more recent recharge:


- High continental latitudes < 4TU
- Low continental latitudes < 1 TU (Fontes, 1979)

Estimating absolute age of young groundwater based on tritium concentration alone is not straight forward. Need assumption for groundwater flow models & mixing with existing groundwater during recharge.

Age commonly can be reliably determined from data on tritium and its decay product (helium-3). The  $3\text{H}/^3\text{He}$  age is based on a calculation that determines the amount of  $^3\text{He}$  derived from radioactive decay of  $^3\text{H}$  in the water.


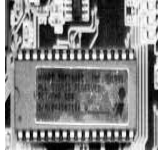

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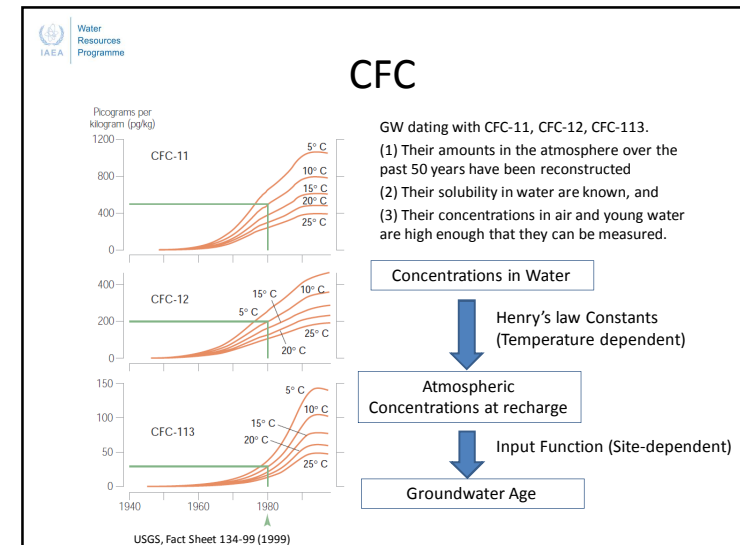
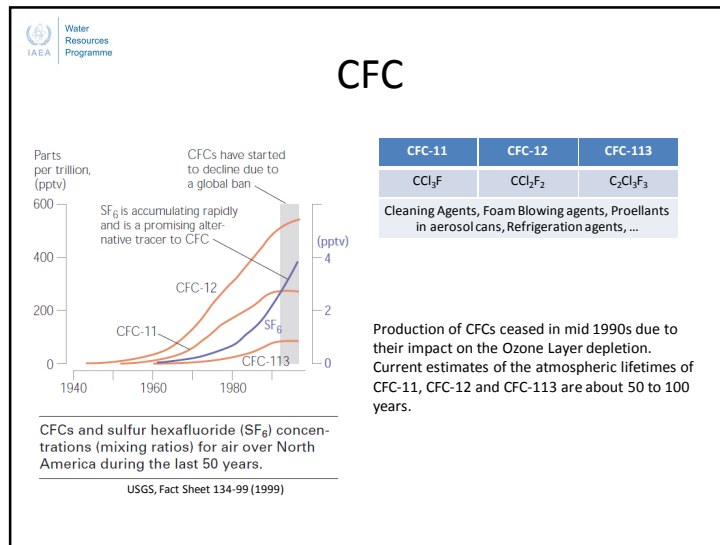
## CFC



CFC-11	CFC-12	CFC-113
$\text{CCl}_3\text{F}$	$\text{CCl}_2\text{F}_2$	$\text{C}_2\text{Cl}_3\text{F}_3$
Cleaning Agents, Foam Blowing agents, Propellants in aerosol cans, Refrigeration agents, ...		

Production of CFCs ceased in mid 1990s due to their impact on the Ozone Layer depletion. Current estimates of the atmospheric lifetimes of CFC-11, CFC-12 and CFC-113 are about 50 to 100 years.

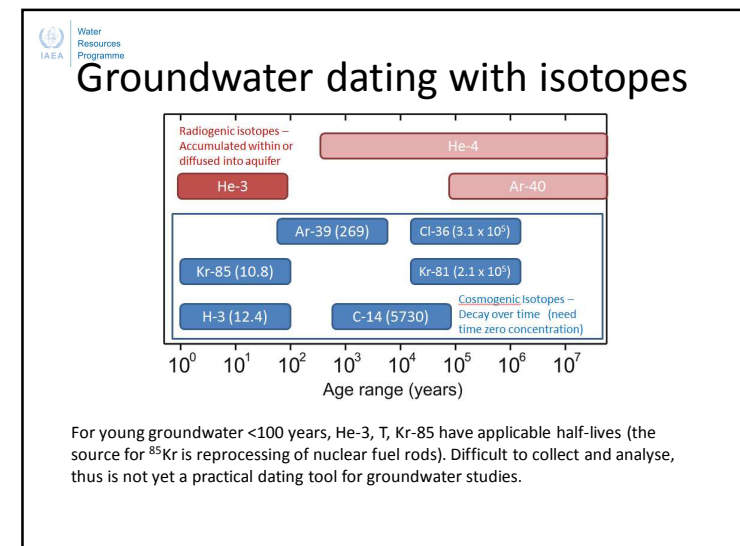






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<sup>4</sup>He & <sup>40</sup>Ar (ages from concentrations and known input rates).



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## Radioactive Decay as a clock

Basic Decay Equations

$$N = N_0 e^{-\lambda t}$$

$$N_0 = N e^{\lambda t}$$

$$D = D_0 + (N_0 - N)$$

$$D = D_0 + N(e^{\lambda t} - 1)$$

D = daughter atoms  
N = parent atoms remaining

By assuming/constraining  $D_0$ , the age can be obtained by present-day concentration of D and N.

The graph shows the percentage of original radioactive isotopes (0 to 100) on the y-axis versus time (0 to 4 half-lives) on the x-axis. A solid blue line represents the parent isotope, which decreases exponentially. A dashed red line represents the daughter isotope, which increases from 0 to 100% over the same period. Vertical dashed lines mark each half-life. Above the graph, a diagram shows a sequence of boxes representing the decay process over four half-lives.

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## Groundwater Dating with $^3\text{H}$ - $^3\text{He}$

- Some characteristics of Helium
  - 5.24 ppm by volume in the atmosphere
  - Not heavy enough to be gravitationally captured in the atmosphere – continuously escaped to space.
  - Escape and addition from inside the earth is now in the steady state with helium residence time of  $10^6$  years.
  - 99.99986% is  $^4\text{He}$  – natural abundance of  $^3\text{He}$  is extremely low (**Sensitive to addition of  $^3\text{He}$  produced by decay of equally rare tritium in water.**)
  - $^3\text{He}/^4\text{He}$  ratios in the air is globally uniform with atmospheric mixing time of a few years. (**Quantification for discriminating  $^3\text{He}$  derived from air and from the tritium decay)**)

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## Groundwater Dating with $^3\text{H}$ - $^3\text{He}$

$$D = D_0 + N(e^{\lambda t} - 1)$$

$$[^3\text{He}]_t = [^3\text{He}]_0 + [^3\text{T}](e^{\lambda t} - 1)$$

$$[^3\text{He}]_0 = \text{constrainable}$$

$$T = \frac{1}{\lambda} \ln\left(\frac{[^3\text{He}]_t}{[^3\text{T}]_t} + 1\right)$$

$^3\text{He}$  at the time of recharge can be estimated from recharge temperature (air contains small amount of  $^3\text{He}$  which will be dissolved into the rain water).

Underlying assumption – The groundwater system should be **closed** in terms of **tritium and its daughter  $^3\text{He}$**  from its isolation from air to sampling.

Called “tritiogenic helium”

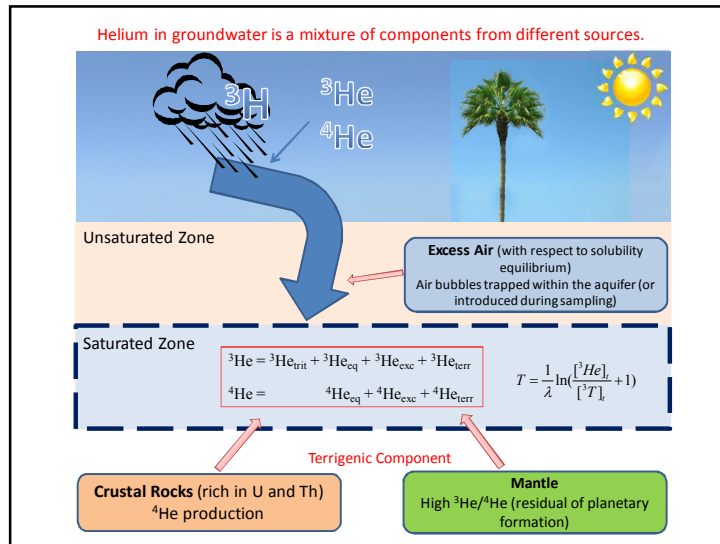
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An ideal case in which fulfils the closed system assumption

The diagram illustrates the process of groundwater recharge. A cloud releases rain containing  $^3\text{H}$  and  $^4\text{He}$ . The rain falls on a palm tree and infiltrates the ground. The ground is divided into an Unsaturated Zone (top) and a Saturated Zone (bottom). In the Unsaturated Zone, equilibrium solubility (p, t, s) controls helium concentration in water unless water is in contact with air – within Vadose Zone (the top of the ground surface to the water table). In the Saturated Zone, the equation  $^3\text{He} = ^3\text{He}_{\text{atm}} + ^3\text{He}_{\text{eq}}$  and  $^4\text{He} = ^4\text{He}_{\text{eq}}$  are shown. The equation  $T = \frac{1}{\lambda} \ln\left(\frac{[^3\text{He}]_t}{[^3\text{T}]_t} + 1\right)$  is also present.

Saturated zone is a closed system with respect to dissolved gases.

**$^3\text{He}$  concentration will begin to rise.**  $^3\text{H}/^3\text{He}$  used to determine a time period for which groundwater has been isolated from the atmosphere (e.g., time since recharge)



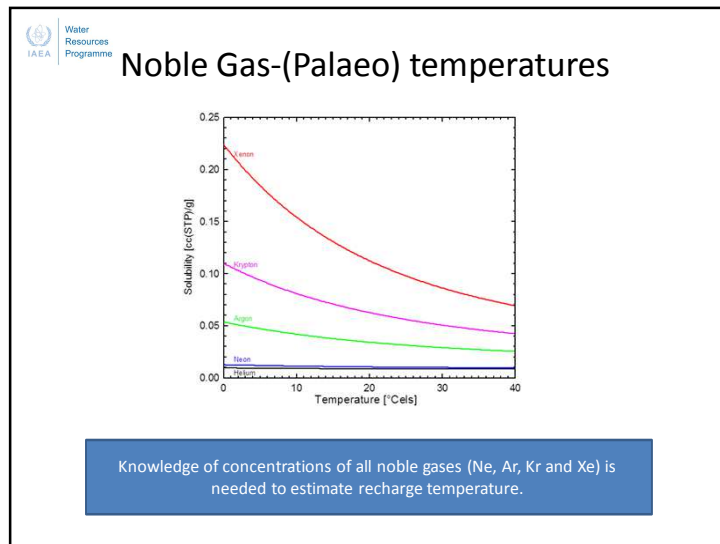
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## T / $^3\text{He}$ Dating of Groundwater: Separation of Helium Components

$$^3\text{He}_{\text{meas}} = ^3\text{He}_{\text{trit}} + ^3\text{He}_{\text{eq}} + ^3\text{He}_{\text{exc}} + ^3\text{He}_{\text{terr}}$$

$$^4\text{He}_{\text{meas}} = ^4\text{He}_{\text{eq}} + ^4\text{He}_{\text{exc}} + ^4\text{He}_{\text{terr}}$$

- $\text{He}_{\text{eq}}$ : Solubility equilibrium, **needs infiltration temperature (noble gases themselves can be a thermometer)**
- $\text{He}_{\text{exc}}$  Excess air determined via Ne
- $\text{He}_{\text{terr}}$  separation possible if either crustal He ( $^3\text{He}/^4\text{He} < 10^{-8}$ ) or mantle He ( $^3\text{He}/^4\text{He} > 10^{-5}$ ) present, not for both



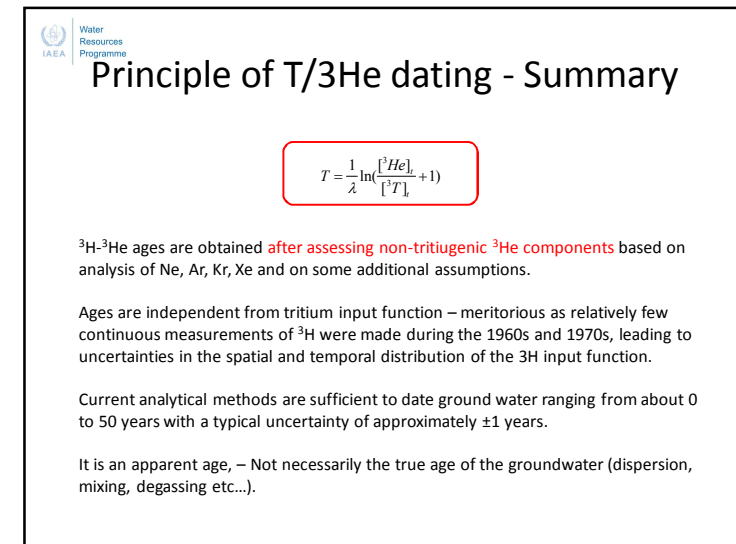
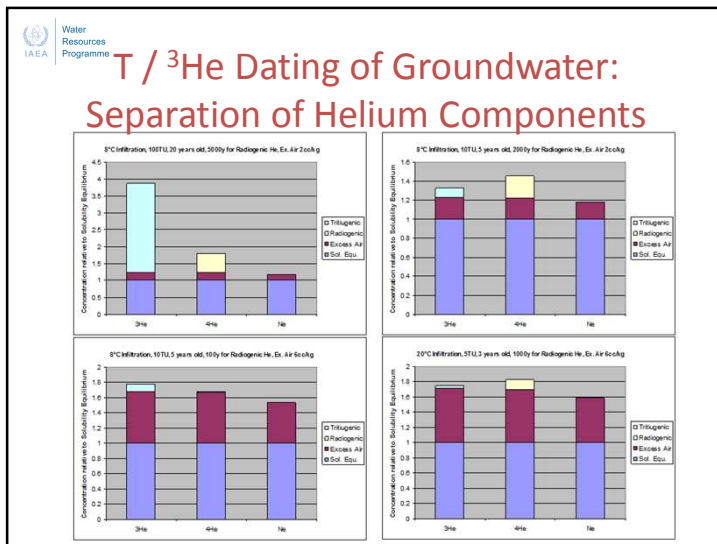
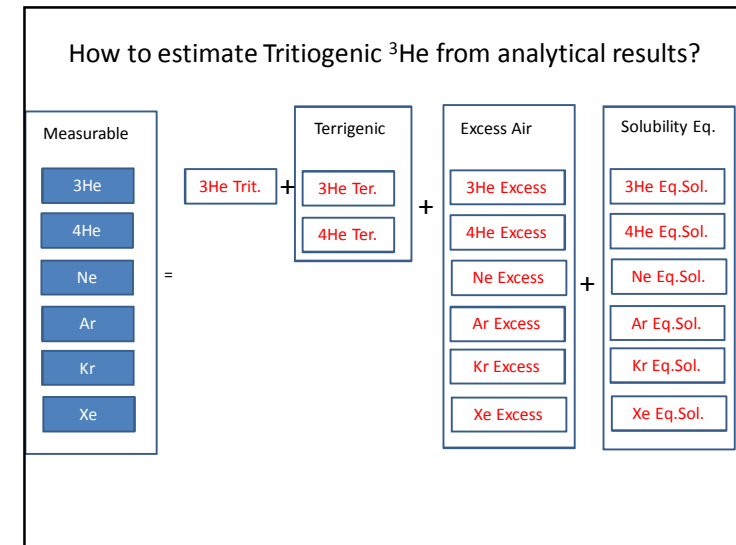
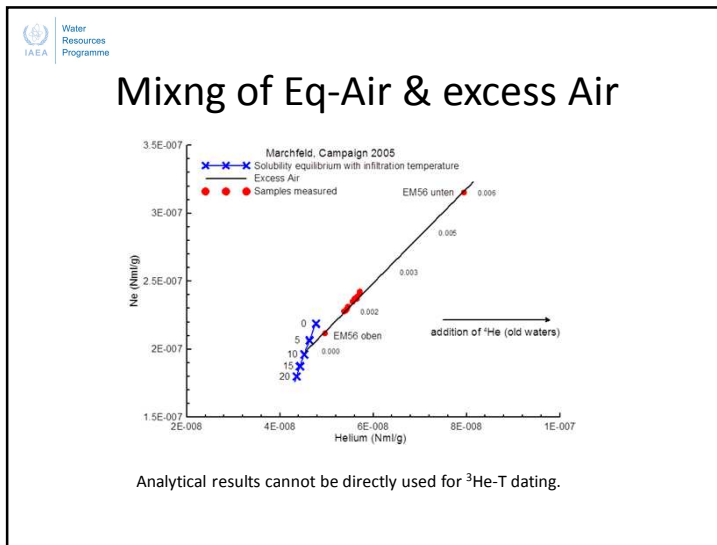
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- $\text{He}_{\text{eq}}$ : Solubility equilibrium, needs infiltration temperature (based on **Ne, Ar, Kr, Xe concentrations**)
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## Measurement Steps for Noble Gases in Water Samples:

**Sampling of water or gas samples in the field**  
Cu-Tube (water sampling), Diffusion/Contactor Samplers (on-site sampling of dissolved gas)



**Gas extraction from water (if Cu-tube is used).**  
Water Vapour Transport through a Capillary Tube  
1 ccSTP in 40g



**Separate Reactive Gases ( $O_2$ ,  $N_2$ ,  $CO_2$ ,  $CH_4$ ...) from Noble Gases**  
Chemical Getters  
Residual: 0.01cc, mainly Argon

**Separate Noble Gases (He, Ne, Ar, Kr, Xe)**  
Cryo techniques down to 10K (-263°C)  
Some amounts: 5e-6cc  $^4He$ , 5e-12cc  $^3He$

**Measure Each Noble Gas in a dedicated Mass Spectrometer**  
High resolution magnetic sector field Mass Spectrometers (SMS) for  $^3He$   
Simple Quadrupole Mass Spectrometers (QMS) for all others



## STEP 1: Sampling



Need to avoid contact with atmosphere during sampling.

Exchange with the atmosphere breaks the closed system principle of the isotope clock.

**Not all wells or pumps are suitable for noble gas sampling**, because of the requirements of closure against air or other gas phases and of sufficient pressure to prevent degassing.

**Closed boreholes are highly desirable**, open wells or springs are problematic, because they allow gas exchange.

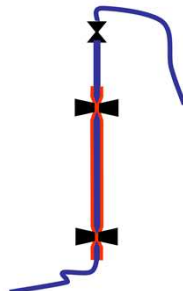
**Ideal are artesian wells with sufficient pressure or wells pumped by a submersible pump.**

Helium is highly diffusive element – They pass through glass/plastic.



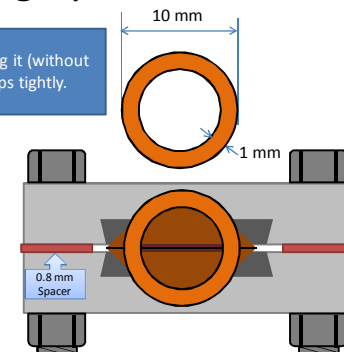
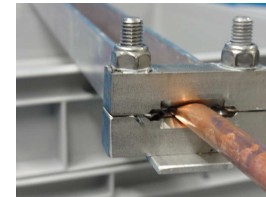
## Water Sampling by Cu-tube

Water is sampled by continuously flowing it (without any bubbles) and by closing two clamps tightly.



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Cu-tube is sealed by "cold-welding" of metal to metal interface.




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## STEP 2: Gas Extraction

Cu-tubes are brought back in the lab and subjected for gas extraction

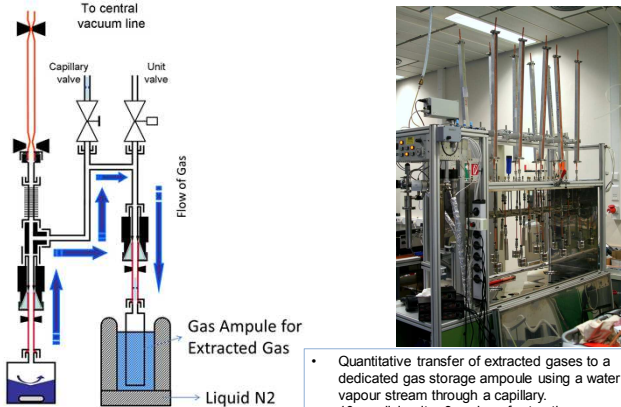
	cm <sup>3</sup> STP	mole
<sup>3</sup> He	3E-12	1E-16
<sup>4</sup> He	2E-06	8E-11
<sup>20</sup> Ne	7E-06	3E-10
<sup>40</sup> Ar	1E-02	6E-07
<sup>84</sup> Kr	2E-06	7E-11
<sup>132</sup> Xe	1E-07	4E-12

+ Major Gases, such as N<sub>2</sub>, O<sub>2</sub>, CO<sub>2</sub>, CH<sub>4</sub> etc....



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## STEP 2: Gas Extraction

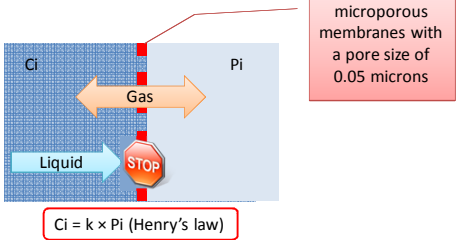


Quantitative transfer of extracted gases to a dedicated gas storage ampoule using a water vapour stream through a capillary.  
10 parallel units, 3 cycles of extraction per week (> 100 samples per month).

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## Gas Sampling with membrane

Direct gas sampling is possible in the field by using a membrane technology



microporous membranes with a pore size of 0.05 microns

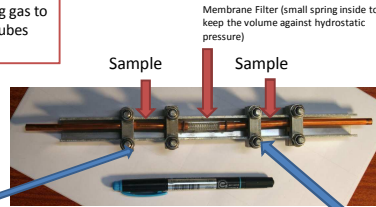
$C_i = k \times P_i$  (Henry's law)

**C<sub>i</sub>:** Concentration of given gas species = important unknown to obtain  
**K:** Henry's constant = known (**water temperature needed**)  
**P<sub>i</sub>:** Partial Pressure of given gas species = will be known parameters by analysis  
 (Molar fraction of noble gas can be determined by mass spectrometer, **Total Pressure of Dissolved Gas needs to be measured in water at sampling**)

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## Gas Sampling with membrane

**Diffusion Sampler:** Sunk into water & allowing gas to diffuse into copper tubes through membrane.

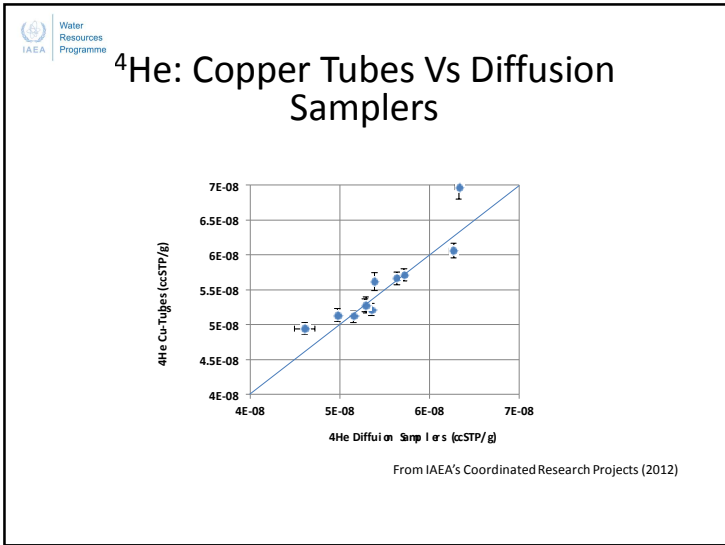
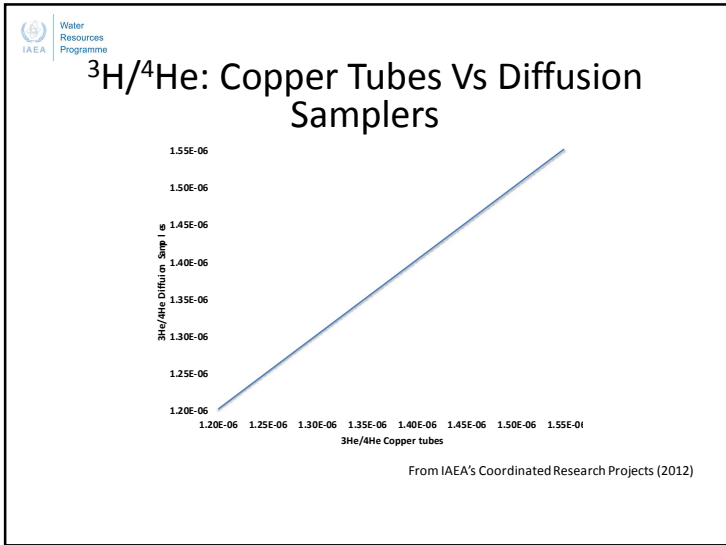


Membrane Filter (small spring inside to keep the volume against hydrostatic pressure)

Outer Clamps (Closed before sinking into water)

Inside Clamps (Closed after at least 24 hours for equilibration)

**Pros:** Light & Small (easy handling), multiple sampling  
**Cons:** Time consuming (>24 hrs of waiting)  
 Need probes for water temperature & Total Dissolved Gas Pressure (TDGP)  
 Time integration over sampling period if water composition changes with time



## Gas Sampling with membrane

**Contactor Sampler:** A compact device that can extract gas from continuous flow of sample water

Shorter extraction time than diffusion samplers.  
No external probes required  
Weights 10 kg – still lighter than a full box of Cu-tubes.

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## Gas Sampling with membrane

**Contactor Sampler:** A compact device that can extract gas from continuous flow of sample water

A watertight, crushproof and dust proof case (85.4 x 32.4 x 17.1 cm).

Water temperature & total Pressure of sampler can be directly measured by a built-in sensors (No temperature & TDGP probes required)

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### Sampling of water or gas samples in the field

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### Gas extraction from water (if Cu-tube is used).

Water Vapour Transport through a Capillary Tube

**1 ccSTP in 40g**



### Separate Reactive Gases ( $O_2$ , $N_2$ , $CO_2$ , $CH_4$ ...) from Noble Gases

Chemical Getters

Residual: 0.01cc, mainly Argon

### Separate Noble Gases (He, Ne, Ar, Kr, Xe)

Cryo techniques down to 10K (-263°C)

Some amounts: 5e-6cc  $^4He$ , 5e-12cc  $^3He$

### Measure Each Noble Gas in a dedicated Mass Spectrometer

High resolution magnetic sector field Mass Spectrometers (SMS) for  $^3He$   
Simple Quadrupole Mass Spectrometers (QMS) for all others



## Noble Gas Measurement



Noble Gas System @ the IHL, IAEA

VG5400 + 3 QMS + Dedicated Gas Handling System

Will be updated to new generation mass spec (Thermo Helix SFT)



## Noble Gas Measurement

Amount of Noble Gases in 1 sample

	cm <sup>3</sup> STP	mole
$^3He$	3E-12	1E-16
$^4He$	2E-06	8E-11
$^{20}Ne$	7E-06	3E-10
$^{40}Ar$	1E-02	6E-07
$^{84}Kr$	2E-06	7E-11
$^{132}Xe$	1E-07	4E-12

We only have a limited amount of noble gases so that the process needs to be done under very clean static vacuum ( $10^{-7}$  ~  $10^{-8}$  mbar)

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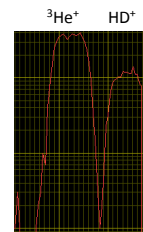


### Sector-type Mass Spectrometer

High Mass Resolution  
necessary to separate  $^3He$  and  $HD^+$  peaks)

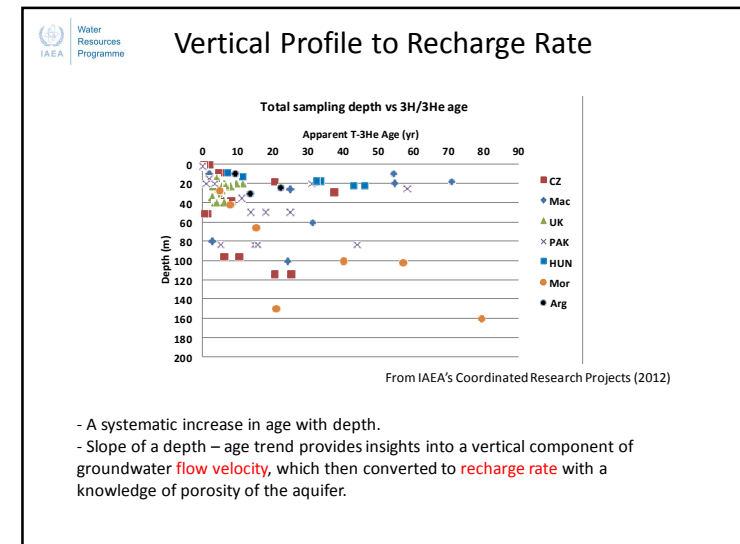
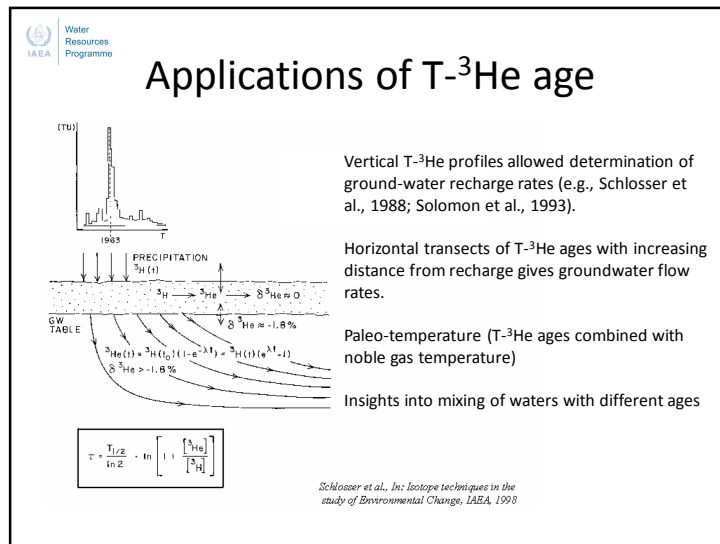
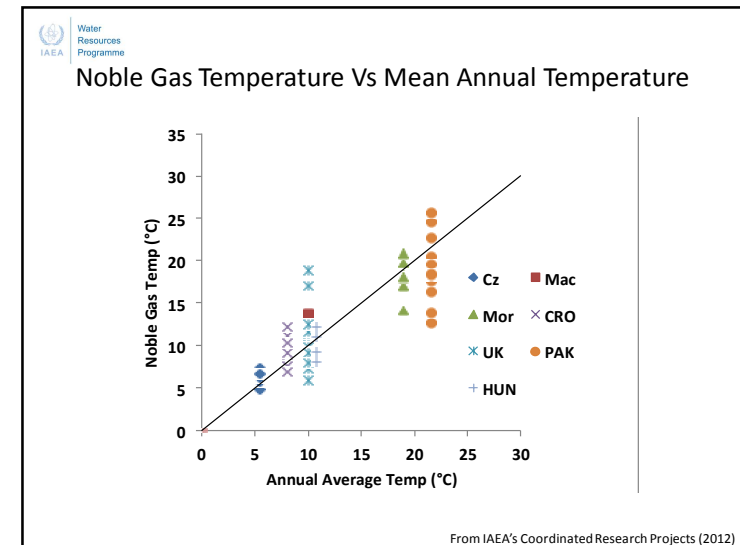
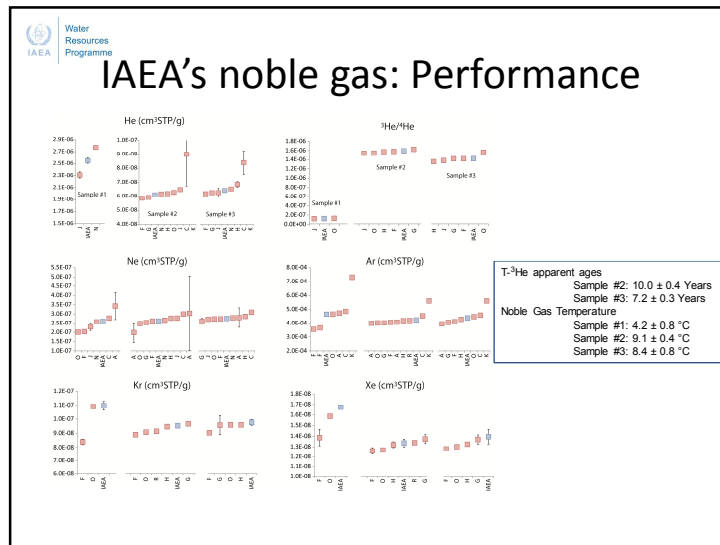
High Sensitivity  
1 cps =  $1 \times 10^{-14}$  cm<sup>3</sup>STP of  $^3He$

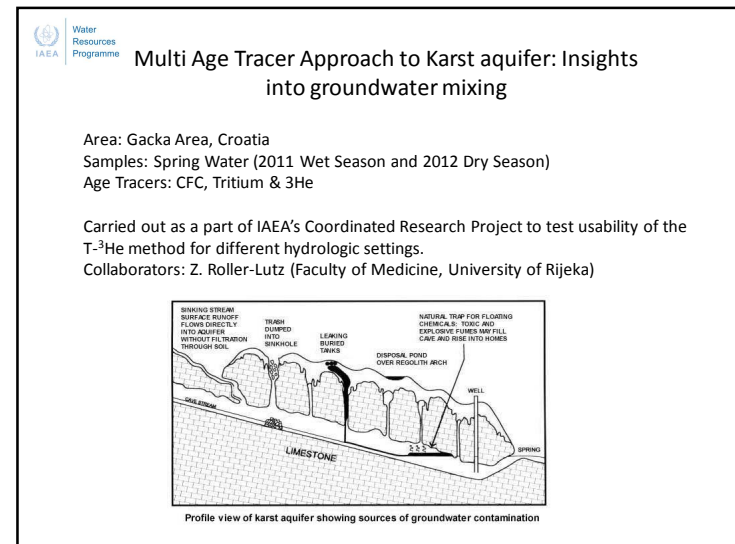
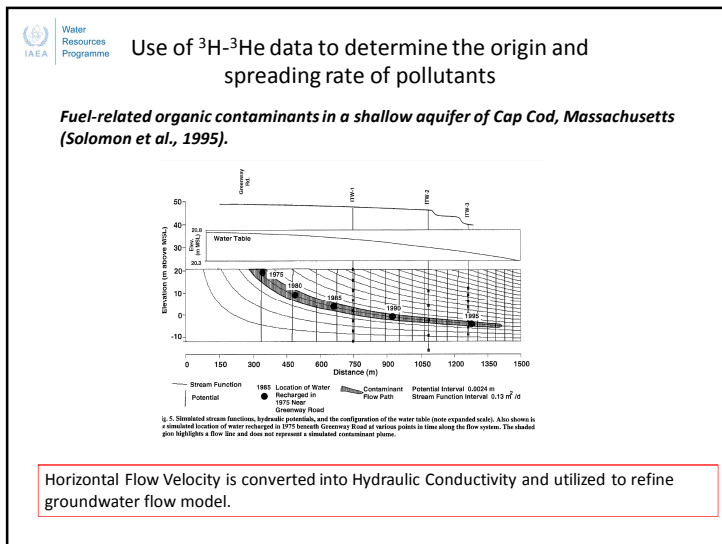
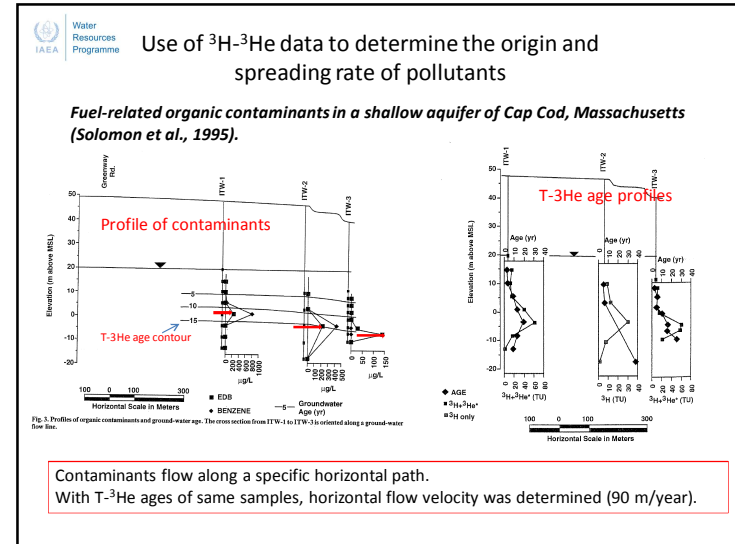
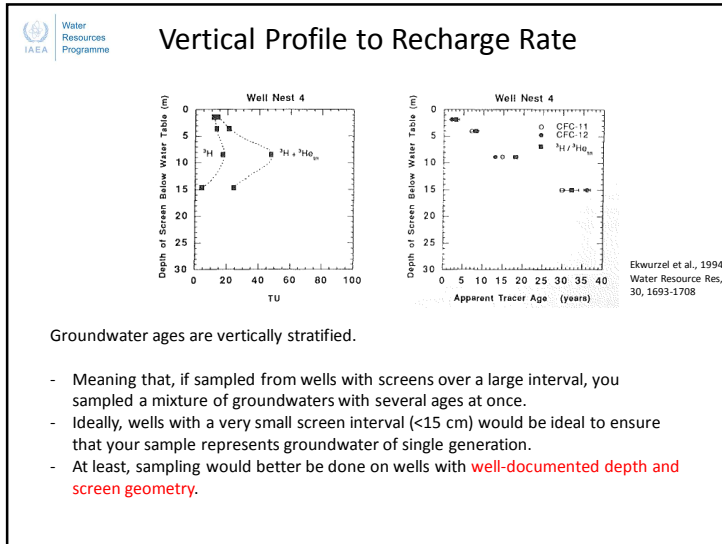
**Dedicated to helium analysis only**

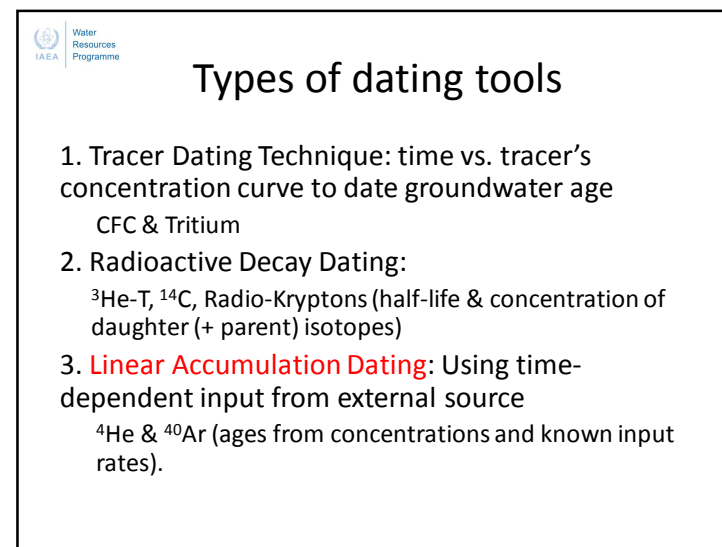
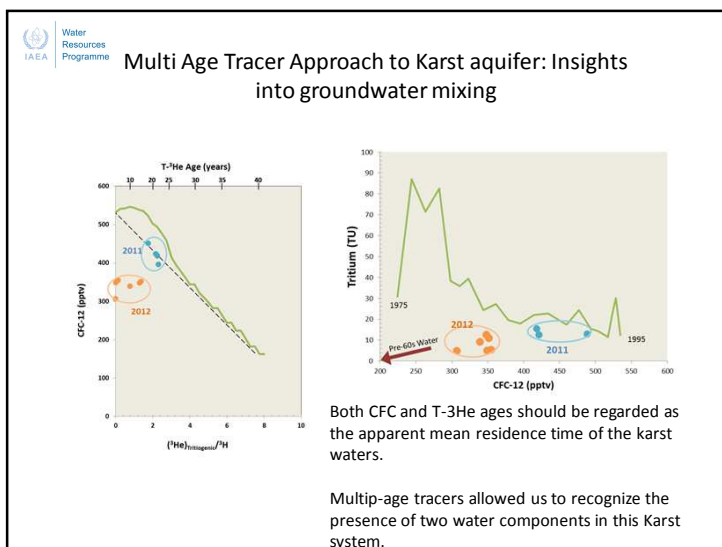
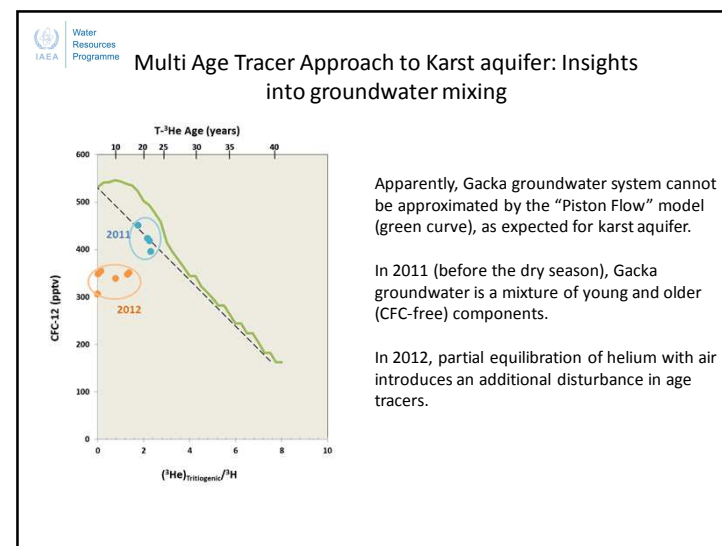
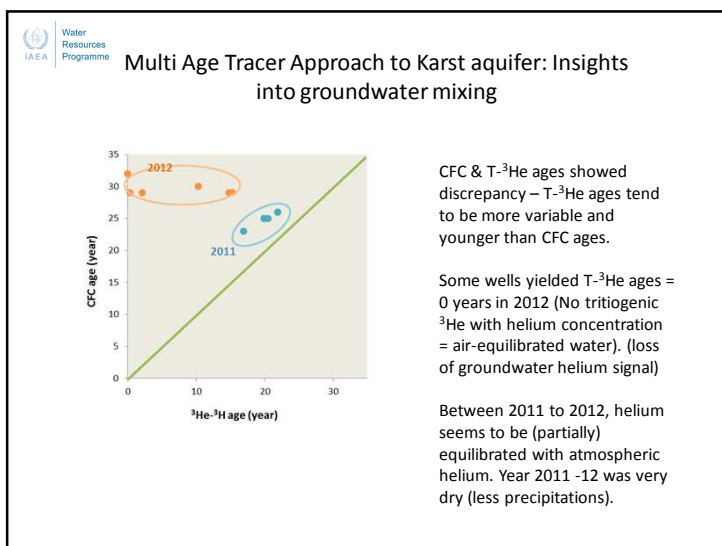


### Measure Each Noble Gas in a dedicated Mass Spectrometer

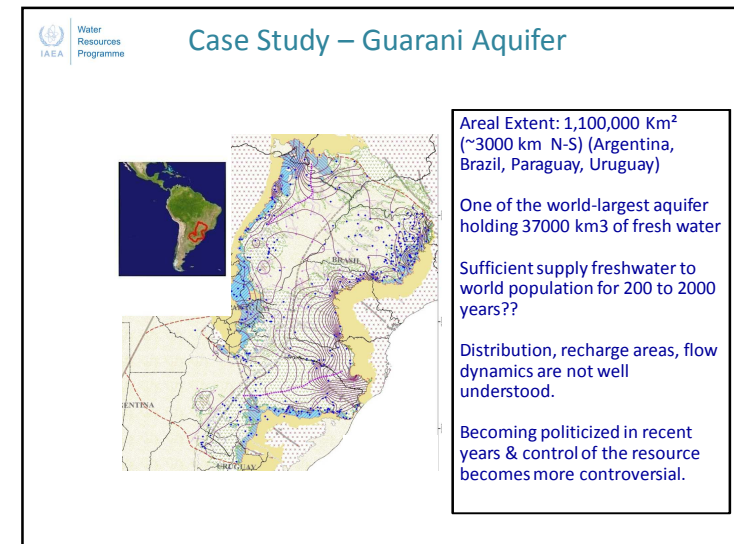
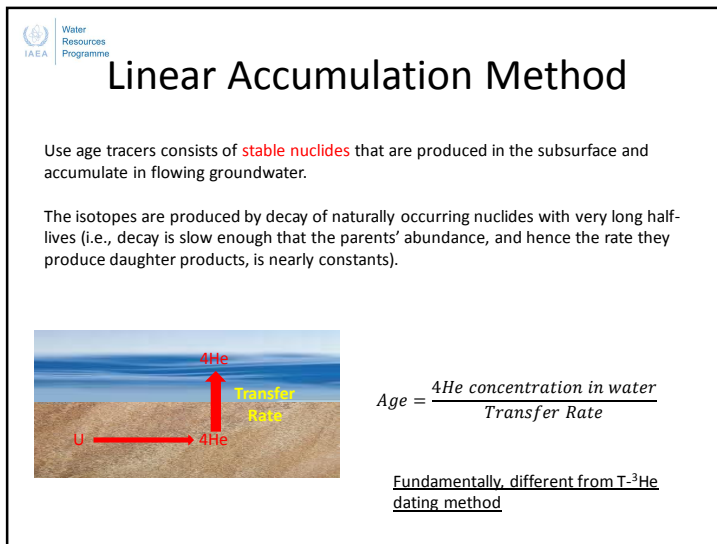
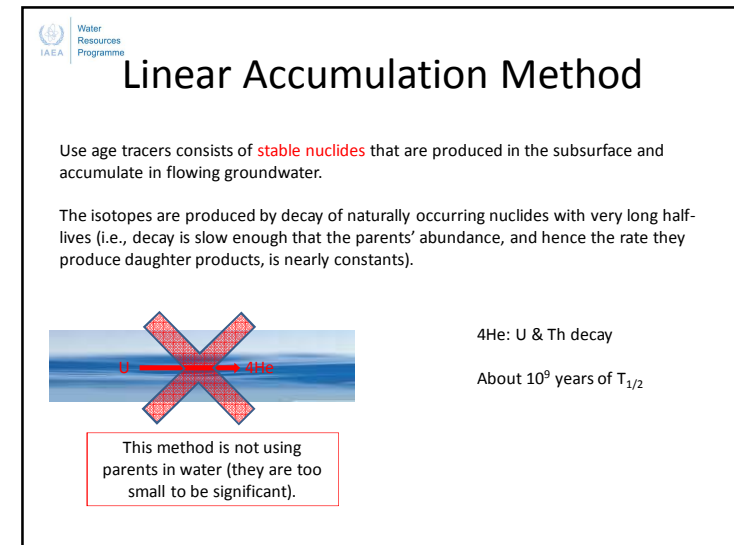
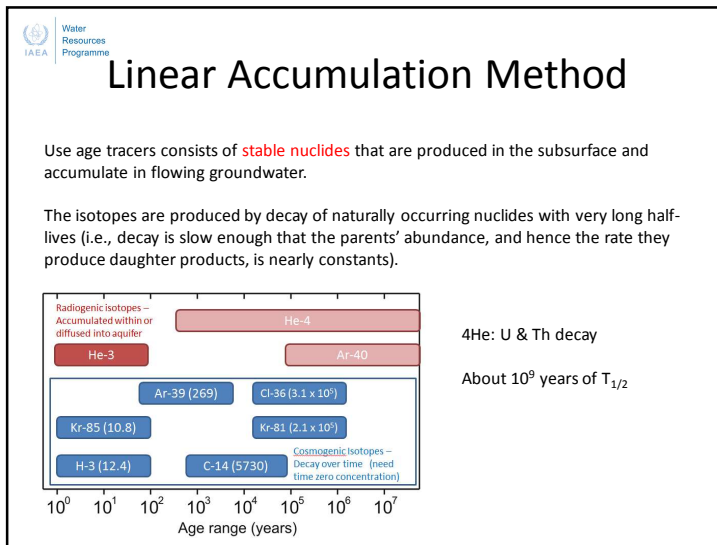
High resolution magnetic sector field Mass Spectrometers (SMS) for  $^3He$   
Simple Quadrupole Mass Spectrometers (QMS) for all others

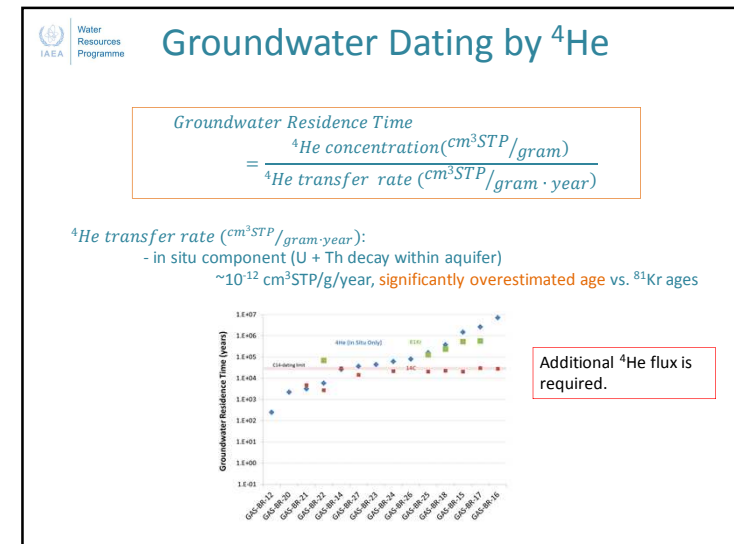
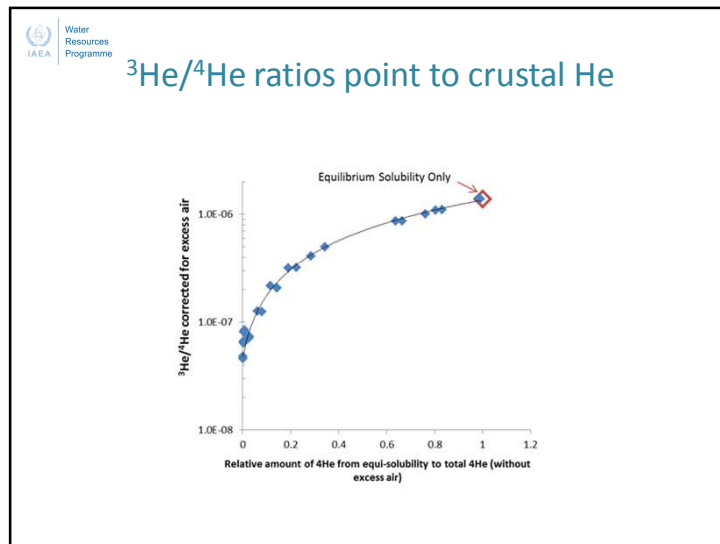
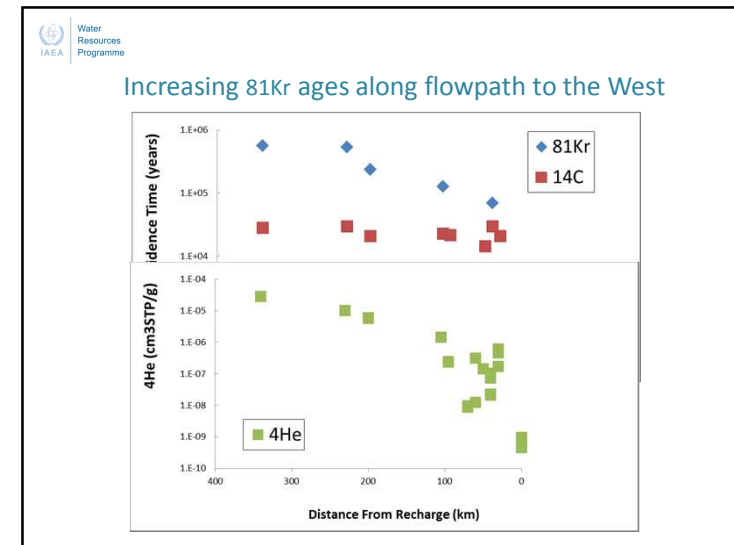
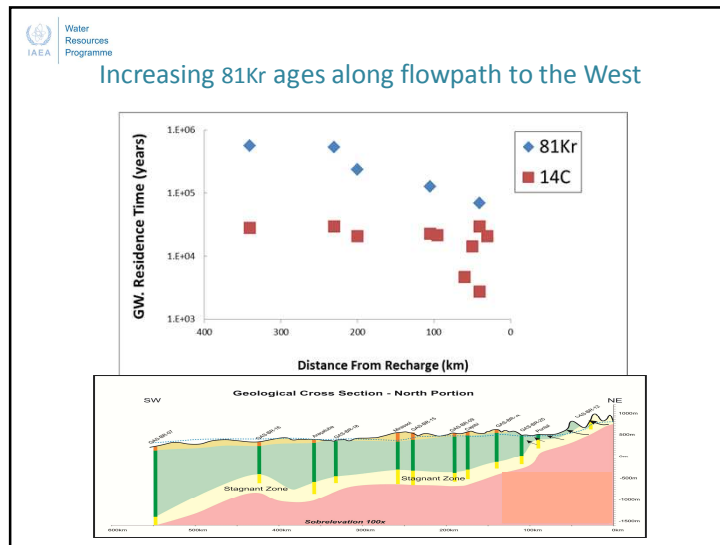




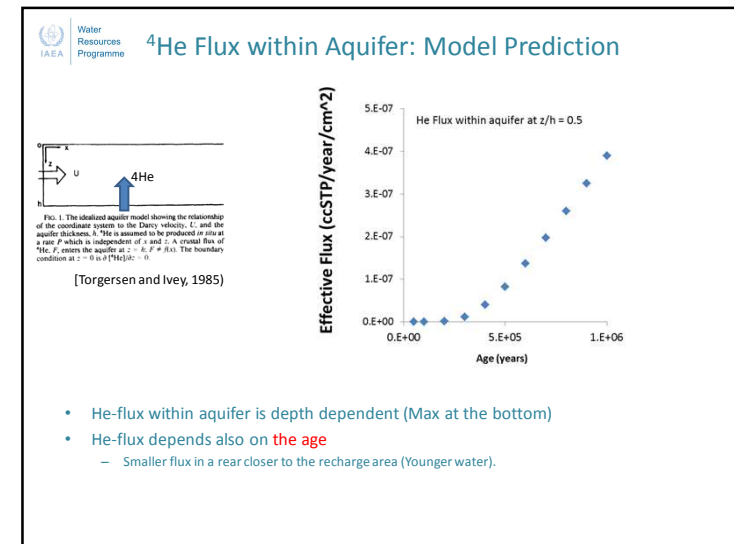
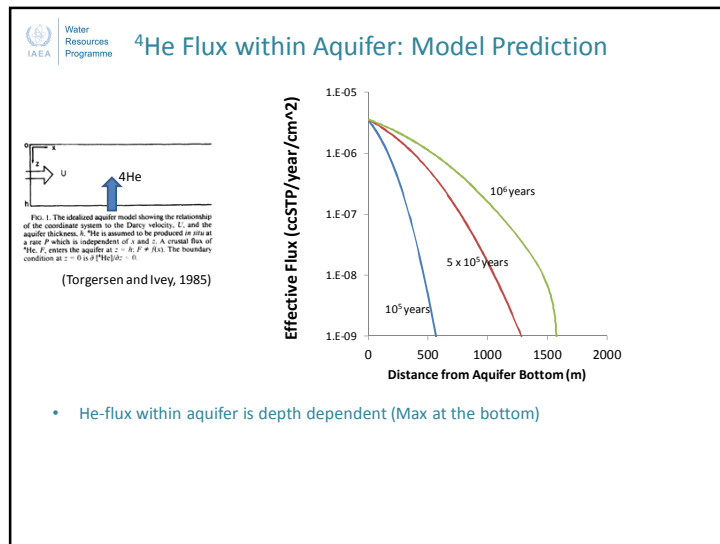
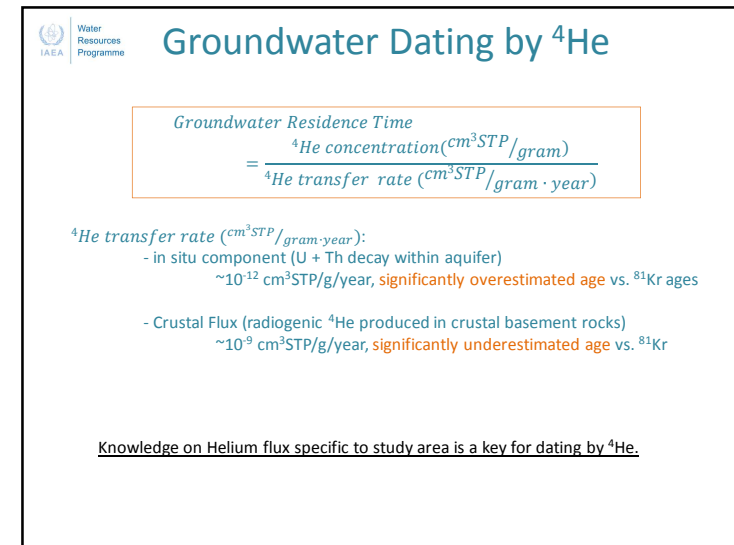
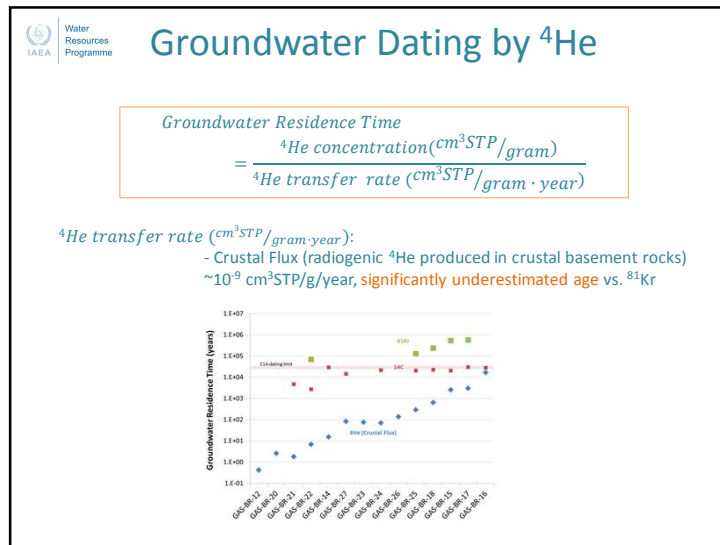












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## Groundwater Dating by $^4\text{He}$

Groundwater Residence Time

$$= \frac{{}^4\text{He concentration (cm}^3\text{STP/gram)}}{{}^4\text{He transfer rate (cm}^3\text{STP/gram} \cdot \text{year)}}$$

$^4\text{He}$  transfer rate (cm<sup>3</sup>STP/gram·year):

- in situ component (U + Th decay within aquifer)  
~10<sup>-12</sup> cm<sup>3</sup>STP/g/year, significantly overestimated age vs.  $^{81}\text{Kr}$  ages
- Crustal Flux (radiogenic  $^4\text{He}$  produced in crustal basement rocks)  
~10<sup>-9</sup> cm<sup>3</sup>STP/g/year, significantly underestimated age vs.  $^{81}\text{Kr}$

Difficulty in obtaining appropriate helium flux is the major obstacle in  $^4\text{He}$  dating.

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## Comparison with GW flow model

FIG. 1. The idealized aquifer model showing the relationship of the coordinate system to the Darcy velocity,  $U$ , and the aquifer thickness,  $h$ .  $^4\text{He}$  is assumed to be produced in situ at a rate  $P$  which is independent of  $x$  and  $z$ . A crustal flux of  $^4\text{He}$ ,  $F$ , enters the aquifer at  $z = 0$ ,  $F = F(x)$ . The boundary condition at  $z = 0$  is  $\partial({}^4\text{He})/\partial z = 0$ .

[Torgersen and Ivey, 1985]

- Observed  $^4\text{He}$  concentrations and  $^{81}\text{Kr}$  age agreed well with groundwater flow model with  $^4\text{He}$  flux from the aquifer base.
- Might be possible to directly estimate groundwater residence time from  $^4\text{He}$  concentration without having a difficulty in parameterizing  $^4\text{He}$  flux distribution (that depends on age and depth).

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- Further works are in progress in the IAEA.

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## Types of dating tools

- Tracer Dating Technique: time vs. tracer's concentration curve to date groundwater age  
CFC & Tritium
- Radioactive Decay Dating:  
 $^3\text{He}$ - $\text{T}$ ,  $^{14}\text{C}$ , Radio-Kryptons (half-life & concentration of daughter (+ parent) isotopes)
- Linear Accumulation Dating: Using time-dependent input from external source  
 $^4\text{He}$  &  $^{40}\text{Ar}$  (ages from concentrations and known input rates).

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## Concluding Remarks

Each dating technique has advantages and limitations – Important to realize these before using outcomes of analysis as “groundwater ages”. Water is mobile, and tracers are not 100 % conservative – Isotope ages are always considered to be model age and/or apparent age.

Combination of information from multiple isotope methods is desirable.

- If agreed, more confidence
- If not, discrepancy itself can be valuable information to identify processes within aquifers.

In any case, selection of appropriate sampling sites and wells, as well as proper sampling is an important first step for groundwater dating.

IAEA Water Resources Programme

## IAEA | Water Resources Programme

### Activities

**Research and development** →

- Coordinated Research Projects (CRP)
- Development/adaptation of field and laboratory methods

**Technology transfer** → Technical Cooperation Projects

**Isotope Hydrology Lab.**

Operation of Global Isotope networks and isotope data dissemination

Analytical services/support to labs in Member States

Quality Control of isotope analysis – Intercomparison Exercises

Education and training

Information exchange (Symposia, workshops, sci. meet.)


Partnerships with other UN-Agencies (UN-Water) and other organizations

IAEA Water Resources Programme

## IAEA | Isotope Hydrology Lab (IHL)

### Facility:

- Noble Gas Mass Spectrometry
  - VG5400
  - Helix 5FT (will be installed in 2013)
- Stable Isotope-Ratio Mass Spectrometry
  - Delta Plus
- Laser Absorption Spectroscopy Analyzer
  - Picarro L1102 x 1 & LGR DLT 100 x 4
- Liquid Scintillation Counters (Tritium)
- CFC's and Radon-222
- Radioisotopes (81Kr, in progress)





IAEA Water Resources Programme


## IAEA | Isotope Hydrology Lab (IHL)

### ACTIVITIES:

**Capacity Building of Member States**

- WR programme is setting up labs in member states
- Training Courses (1-2/year on laser Absorption Analyzer)
- Lecturing courses held in member states
- Day to day supports on external labs



Water  
Resources  
Programme

IAEA | Isotope Hydrology Lab (IHL)


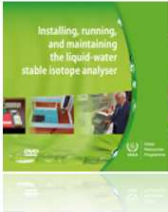
ACTIVITIES:

Education/Training Materials, Software


Training Courses (2/year on laser Absorption Analyzer)

E-learning

Day to day supports on external labs



LIMS for Lasers



The IAEA Isotope Hydrology Laboratory (IHL) has partnered with the U.S. Geological Survey to develop a new Laboratory Information Management System (LIMS) for water oxygen and hydrogen isotope analysts for users of laser absorption spectrometers. LIMS for Lasers is a user-friendly MS Access-based instrument and laboratory data management system for all off-line (LA-ICP-MS) or on-line (mg-down (CROD) laser absorption spectrometers (LAS) used for  $\delta^{18}\text{O}$  and  $\delta^2\text{H}$  assays of water. This open-source database software will assist new and current users of LAS instruments to improve their laboratory productivity and QA/QC outcomes, and is available cost-free for all laboratories.

