

Models coupling with specific focus on ocean physics and biogeochemistry

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OUTLINE

- ✓ Model coupling
- ✓ What is biogeochemical modelling
- ✓ The OFFLINE coupling
- ✓ Some technical issues



The need for model coupling



http://www.esa.int/spaceinimages/Images/2006/09/The_Changing_Earth

- > Atmosphere (air)
- Hydrosphere (groundwater, glaciers, oceans, lakes and streams)
- Cryosphere (sea ice, glaciers, snow)
- Biosphere (sum of all ecosystems, presence of DNA)
- Pedosphere (soil)
- Lithosphere (crust and uppermost solid mantle)
- Anthroposphere (human activities)

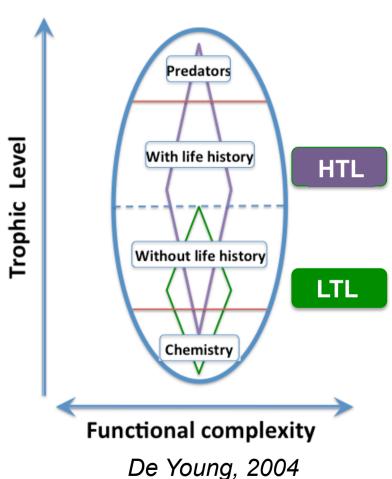
"COUPLING" → flow of mass, momentum, energy between spheres

... In the marine biosphere itself ...

Many modules/approaches are present, each related to different scientific communities

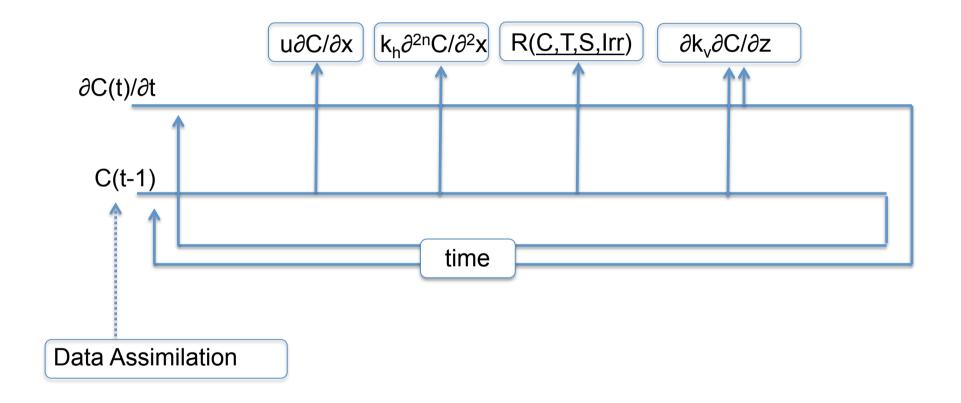
- Plankton organisms (Lower Trophic Levels)
- Fishes (Higher Trophic Levels)
- Benthic/Sediment models ("the pedosphere
 @ the ocean bottom")

"COUPLING" → flow of mass due to predation dynamics or sedimentation/resuspension



... Coupling in terms of forcings ...

Forcings coming from other models (e.g. circulation in a transport model), or model trajectory correction (e.g. 3DVAR)



"COUPLING" → forcing from an external model on a specific operator e.g. advection



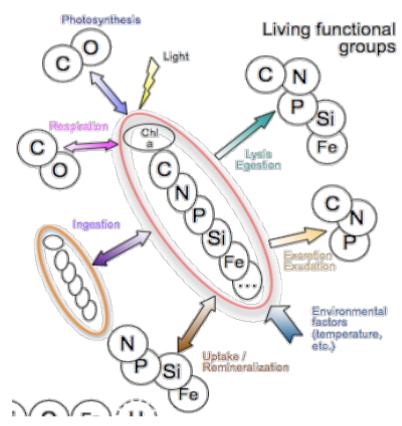
What is marine biogeochemistry?

This talk → focus on marine biogeochemistry

It is about the distributions of chemical concentrations in the Sea and the processes that control them

What is the mean abundance of key elements

What is the spatial and temporal variance of such distributions, what control such variability?



Vichi et al., 2007

Spatial scales: from global ocean to coastal areas

Temporal scales: hindcast mode, forecast, scenario projection



Basic equations

Eulerian framework, conservation equations for chemical concentrations *C*

$$\frac{\partial C}{\partial t} = \left. \frac{\partial C}{\partial t} \right|_{advection} + \left. \frac{\partial C}{\partial t} \right|_{diffusion} + R(C, T, S, Irr, \dots)$$

R → reaction terms (Source minus Sink)

$$\frac{\partial C}{\partial t} = -\mathbf{U} \cdot \nabla C + \nabla (\mathbf{D} \nabla C) + R(\mathbf{C}, T, S, Irr, \dots)$$

"COUPLING" →OGCM

(Navier Stokes equation solver)

es. NEMO MITgcm ROMS

"COUPLING" → REACTOR

(Biogeochemical fluxes)

es. BFM ERSEM

. . .



Biogeochemical Reactors

R

Uptake

Release

Predation

Organic Matter

Semi-labile DOM

Semi-refract. DOM

Particulate detritus

Labile DOM

Predation

Respiration

Release

Inorganic

Carbon dioxide

Reduction Equiv.

Inorg. Nutrients

Phosphate

Ammonium

Bioav. Iron

Nitrate

Silicate

Respiration

Oxygen

Alkalinity

Zooplankton

Carnivorous Meso

Omnivorous Meso

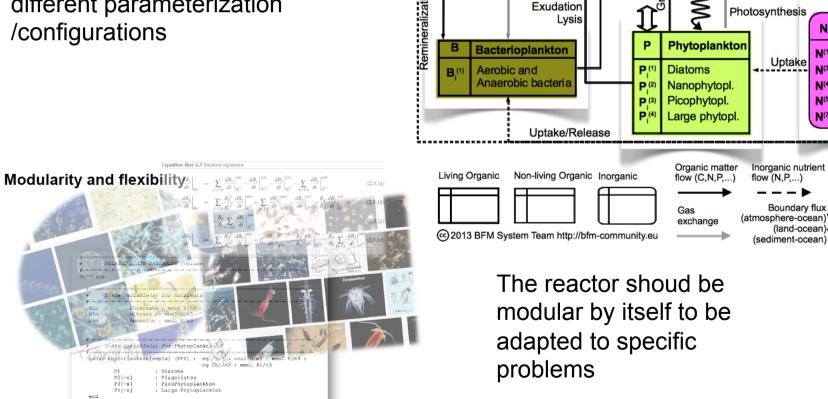
Microzooplankton

Heterotrophic nanoflagellates

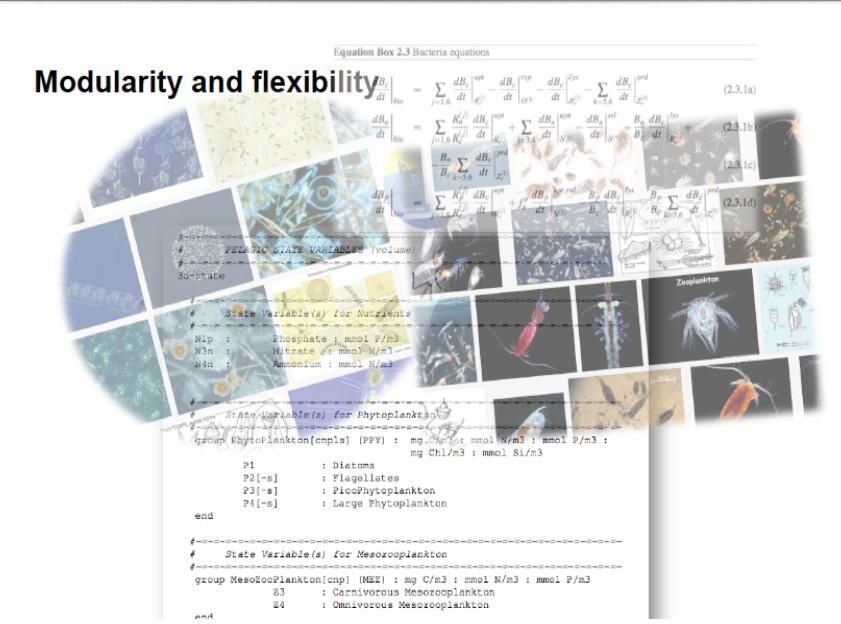
Biogeochemical modelling is usually a problem oriented discipline many different models with different parameterization /configurations

State Variable(s) for Mesozopplankton

group MesoZooPlankton[cnp] (MEZ) : mg C/m3 : mmol N/m3 : mmol P/m3



Biogeochemical Reactors





... Some examples ...

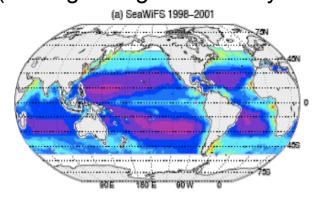


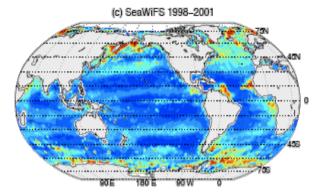
Global Ocean applications

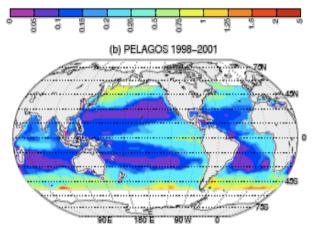
- ✓ BFM coupling with the GCM NEMO at 0.25° and 2°
- ✓ Hindcast simulations of the global ocean biogeochemistry (Vichi et al. 2007a,b;
 Vichi and Masina, 2009)
- ✓ Biogeochemical cycles in the Earth System under current and future climate conditions with the CMCC Earth System Model (Vichi et al., 2011; Patara et al., 2012)

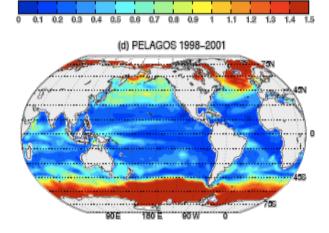
PELAGOS

(PELAgic biogeochemistry for Global Ocean Simulations)











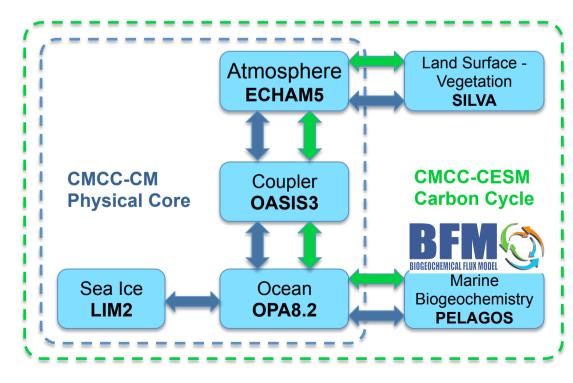


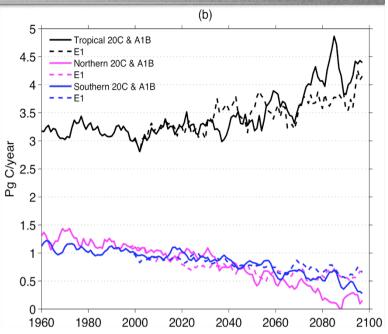




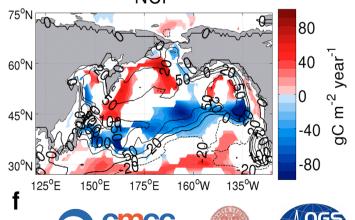
Earth System Modelling

PELAGOS is the marine biogeochemistry component of the CMCC-CESM Carbon Earth System Model that participated to the Climate Model Intercomparison Project Phase 5 (Cagnazzo et al., 2013)





Scenario changes of Net Community Production on the Pacific NCP 1990-2000



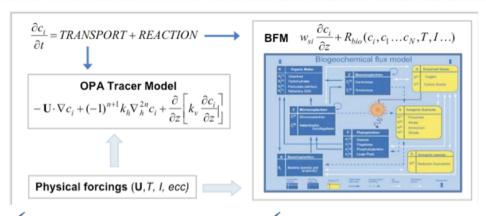








BFM in the Mediterranean Med



Offline approach: coupling with precomputed physical fields from OGCM

- Horiz. Res. = 1/8°
- Vert. Res. = 43/72 levels
- Time Res. = 1800 s

1 year simulated in 2 hours

Model description and forcings

CIRCULATION MODEL

OGCM: NEMO (OPA 9) (Madec 2008

http:/www.nemo-ocean.eu Oddo et al. 2009)

Horizontal Resolution: 1/16 deg ~ 7 Km

Vertical Resolution: Z-coordinates, 71 levels (partial steps)

-Free run: no relaxation to climatology, full freshwater flux

(major rivers), no heat flux correction Parallel simulations (on-line)

BIOGEOCHEMICAL FLUX MODEL

- -BFM: Biological Flux Model (Vichi et al. a,b 2007)
- -Carbon based multi-nutrient food web description Carbon, Nitrogen, Phosphorus and Silica cycles
- -Potential for multiple nutrient co-limitation (Nitrogen, Phosphorus and Silica)
- phytoplancton
- mesozooplancton
- microzooplancton
- bacteria

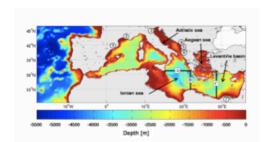


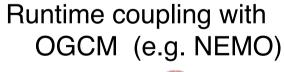
-Physical model settings: ECMWF ERA40 atmospheric forcing functions

-Initial conditions for nutrients and oxygen: annual OA climatologies from SEADATANET project (http://www.seadatanet.org) merged with World Ocean Atlas climatology in the Atlantic

-Initial conditions for biology: homogeneous guesstimates with vertically-distributed analytical

-Nutrient River input: data from Sesame Project (Ludwig et al. 2009)



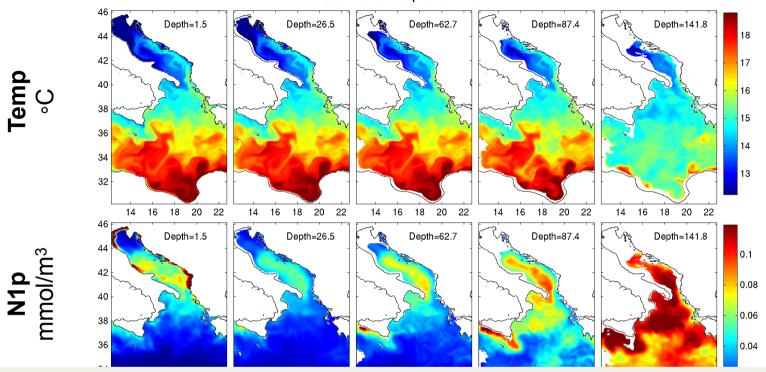


Online approach:



Test case: Adriatic sea + Ionian Sea

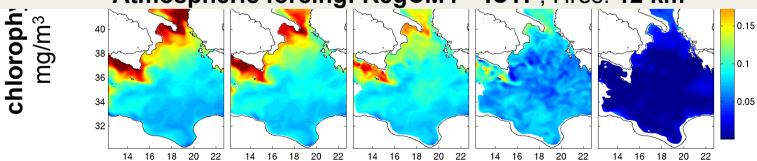
Winter conditions:



Resolution→ H: 1/32°, V: 72 levels

Boundaries: MyOcean Med forecast system, discharges and tracers concentration at rivers, atm. pCO2.







... Different approaches ...

INTERNAL COUPLING

Elements are in the **same** program unit

ON-LINE

Data exchange through working memory (es .mod)

DIRECT

EXTERNAL COUPLING

Elements are in **different** program units e.g. both can be run indipendently

OFF-LINE

Date exchange through external storage (files)
Usually one way

ON-LINE

Date exchange through working memory

DIRECT

INDIRECT Use of a coupler

The best approach depends on the specific problem

Joeckel, 2012 Kerweg & Joeckel, GMD,2012

OFFLINE approach

RE

TM+
REACTION

Applications

- ✓ intermediate processing
- ✓ Operational chain
- √ Climate simulations
- √ Sensitivity analisys

ONLINE approach



OGCM



TM+
REACTION

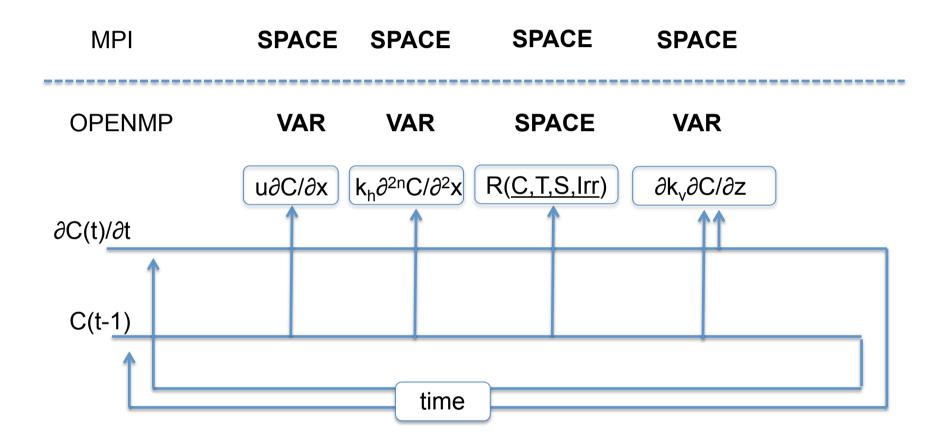
- ✓ MITgcm
- ✓ Nemo



... The OFFLINE APPROACH ...

OFFLINE - The transport component

Different approaches for parallellisation (SPACE → domain decomposition, VAR → parallelisation on biogeochemical variables)



Usually > 24 cores for a real applications



- ✓ To carry out those computations a lot of computational resources are required
- ✓ Attention to optimization in order to carry out experiments in an acceptable time, and reducing I/O problems
- ✓ use of HPC facilities

OFFLINE approach

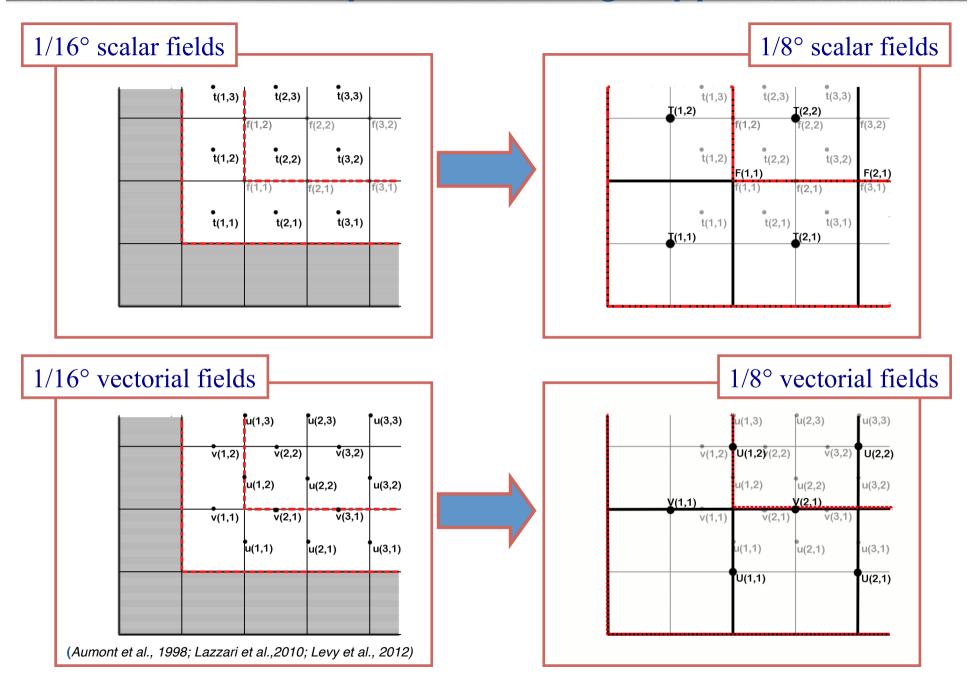


Applications

- ✓ intermediate processing
- ✓ Operational chain
- √ Climate simulations
- √ Sensitivity analisys



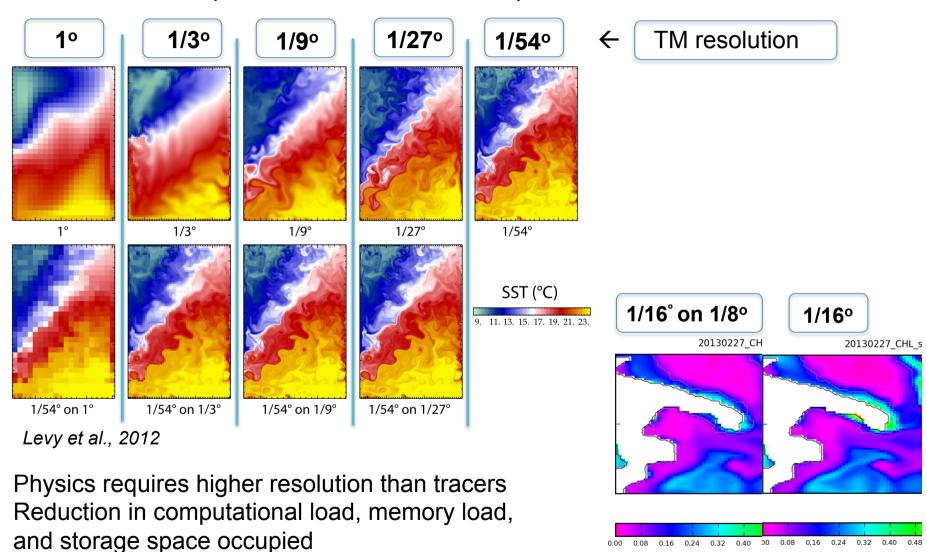
INTER-proc: Cell merge approach





Reduction in computational load

Resolution of the trasport OFFLINE model for temperature



Coastal areas could be affected

OFFLINE approach

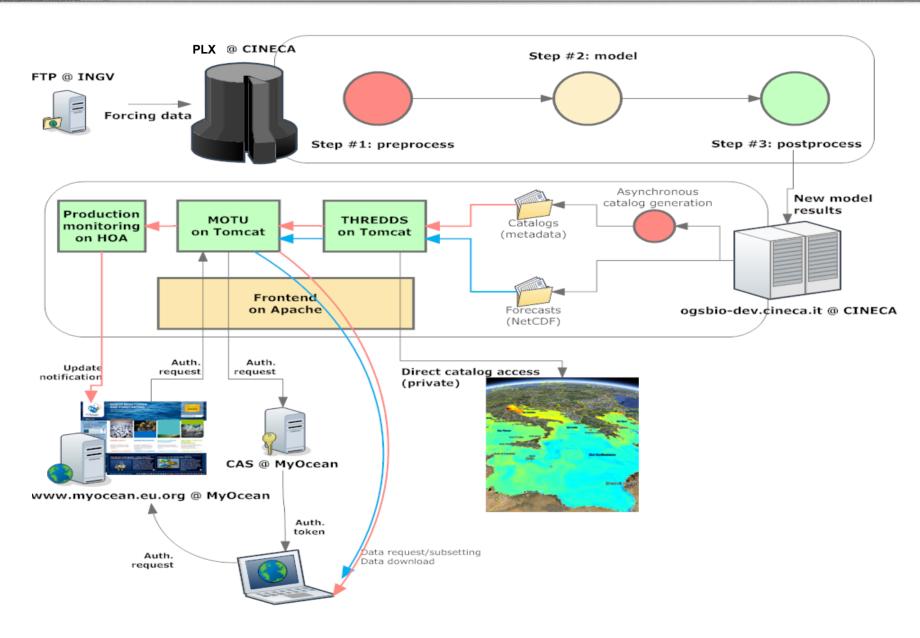


Applications

- ✓ intermediate processing
- ✓ Operational chain
- √ Climate simulations
- √ Sensitivity analisys



Operational Chain in OFFLINE mode





The operational chain



7 days of hindcast/analysis (using INGV physical forcing analysis and ICs via DA based on GOS-ISAC-CNR satellite chlorophyll)



10 days of forecast (using INGV physical forcing forecast)

Initial Conditions from previous run

Saturday run

T F S S M T W T F S S M T W T F S S M T

3DVAR scheme using OC TAC data

Wednesday run

MTWTFSSMTWTFSSMTWTFS

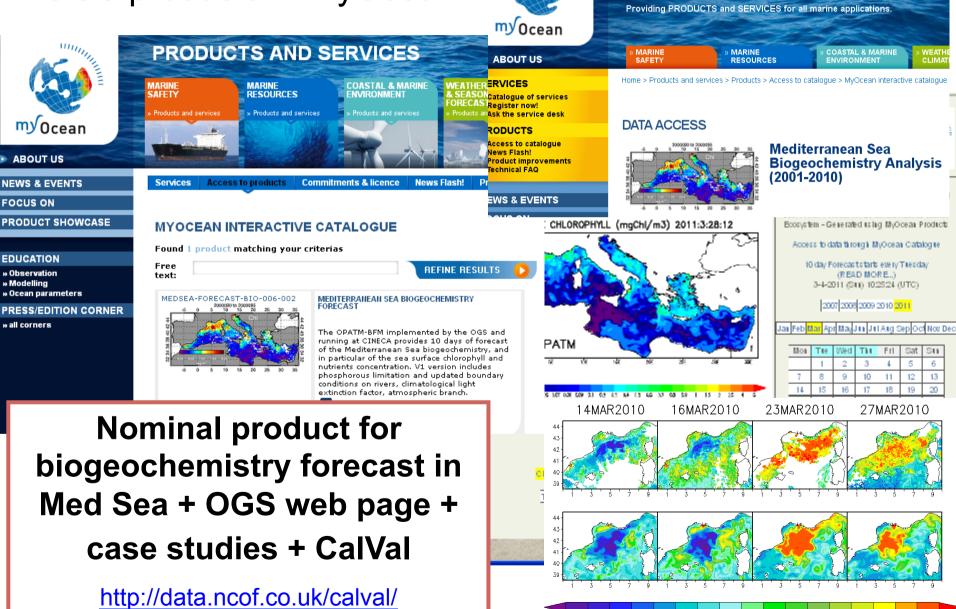
F|S|S|M|T|W|T|F|S|S|M|T|W|T|

Initial Conditions from previous run

Saturday run



OGS products in MyOcean



OCEAN MONITORING

and FORECASTING

0.020.040.060.080.10.20.30.40.50.60.70.80.9

OFFLINE approach



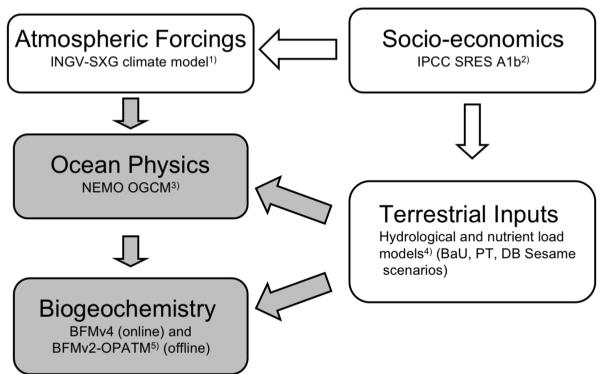
Applications

- ✓ intermediate processing
- ✓ Operational chain
- √Climate simulations
- √ Sensitivity analisys



Simulations on climatic time scales





1) Gualdi et al. (2008); 2) Nakicenovic and Swart (2000); 3) Oddo et al (2009); 4) Ludwig et al. (2010); 5) Lazzari et al. (2014)





- Increase of carbon rates both production (GPP) and community respiration (RSP)
- ✓ Increase of dissolved semi-labile carbon
- ✓ Reduction in biomass

MEDITERRANEAN BASIN

	20C	A1B-BaU	A1B-PT	A1B-DB
GPP	0.66	0.044	0.047	0.029
RSP	0.65	0.044	0.048	0.030
NPP	0.36	0.032	0.036	0.015
NCP	0.01	-0.001	-0.033	-0.064
DSL	0.96	0.038	0.044	0.035
BIO	4.12	-0.046	-0.043	-0.056

WESTERN BASIN

	20C	A1B-BaU	A1B-PT	A1B-DB
GPP	0.81	0.023	0.017	0.011
RSP	0.80	0.023	0.016	0.013
NPP	0.46	0.009	0.001	-0.006
NCP	0.01	0.050	0.019	-0.002
DSL	1.02	0.030	0.027	0.025
BIO	4.97	-0.070	-0.074	-0.076

EASTERN BASIN

	20C	A1B-BaU	A1B-PT	A1B-DB
GPP	0.58	0.061	0.073	0.044
RSP	0.56	0.063	0.076	0.046
NPP	0.30	0.053	0.067	0.034
NCP	0.01	-0.035	-0.068	-0.104
DSL	0.93	0.045	0.056	0.042
ВЮ	3.63	-0.027	-0.018	-0.039

OFFLINE approach

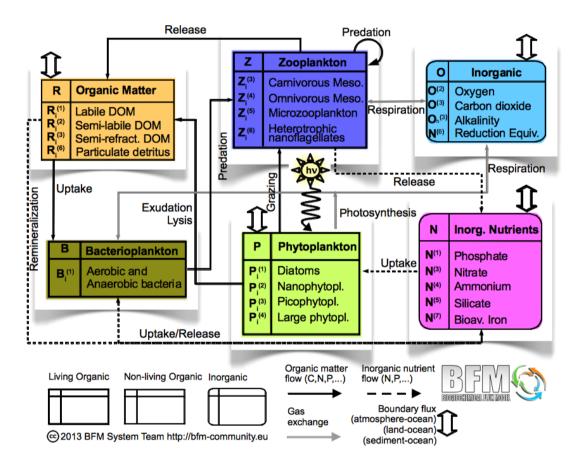


Applications

- ✓ intermediate processing
- ✓ Operational chain
- ✓ Climate simulations
- **✓** Sensitivity analisys



Sensitivity to parameters uncertainty



9 PFT with ~ **20** physiological processes each

~200 parameters control the kinetic of the processes

What is the uncertainty in the model results?

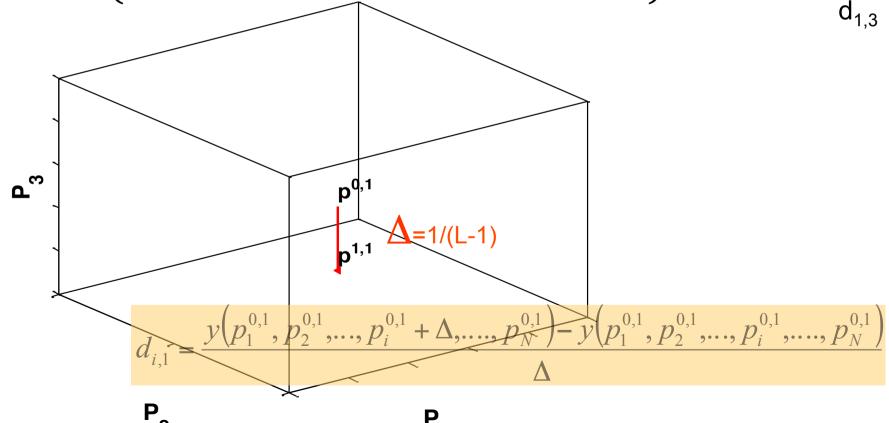
Morris's method

a base vector is randomly chosen within the parameter space

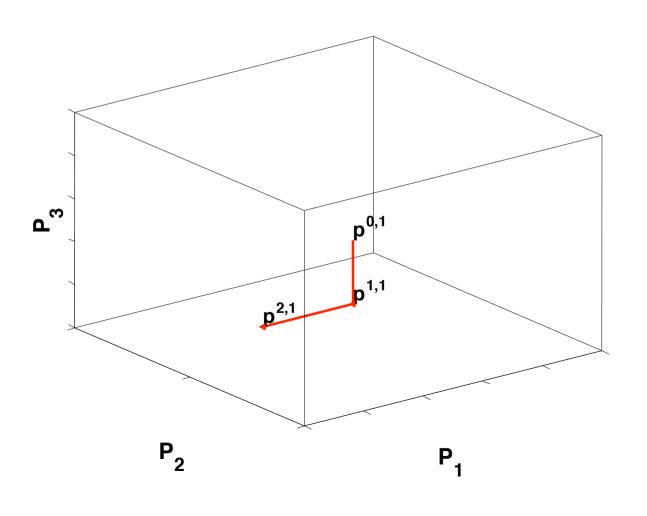
$$\mathbf{P}^{0,1} = \left\{ P_1^{0,1}, P_2^{0,1}, \dots, P_i^{0,1}, \dots, P_N^{0,1} \right\}$$

one parameter is randomly chosen and varied of Δ

elementary effects on $\mathbf{P}^{1,1} = \left\{ P_1^{0,1}, P_2^{0,1}, \dots, P_i^{0,1} + \Delta, \dots, P_N^{0,1} \right\}$ model output



a second parameter is randomly chosen to be varied of Δ

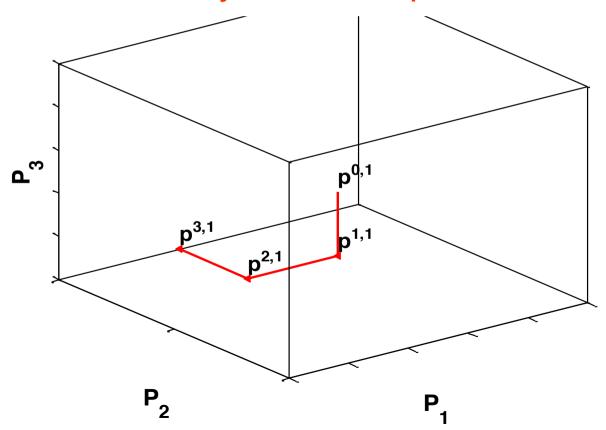


Elementary effects on model output

 $d_{1,1}$ $d_{1,3}$

the step is repeated until all of the parameters have been varied a trajectory on the space of the parameter is then performed

one elementary effect for each param.



 $d_{1,1}$ $d_{1,2}$ $d_{1,3}$

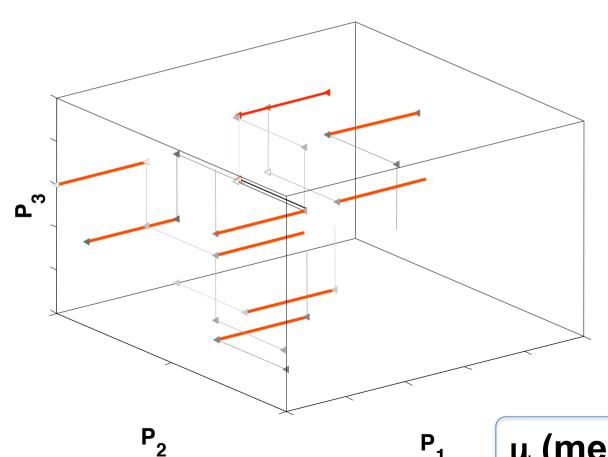


How it works ...3

distribution of elem. effects on the model output due to variations of the first parameter

elementary effects on model

output



$d_{1,1}$	d _{1,2}	$d_{1,3}$
$d_{2,1}$	$d_{2,2}$	$d_{2,3}$
$d_{3,1}$	$d_{3,2}$	$d_{3,3}$
$d_{4,1}$	$d_{4,2}$	$d_{4,3}$
$d_{5,1}$	d _{5,2}	$d_{5,3}$
$d_{6,1}$	d _{6,2}	$d_{6,3}$
$d_{7,1}$	d _{7,2}	$d_{7,3}$
$d_{9,1}$	$d_{9,2}$	$d_{9,3}$
$d_{9,1}$	$d_{9,2}$	$d_{9,3}$
$d_{10,1}$	d _{10,2}	d _{10,3}

P₁

 μ (mean), σ (st.dev.)

To carry out this experiment ...

... on a 3D implementation of BFM

- ✓ sensititity of 165 parameters on 122 variables in 51 different conditions (trajectories)
- ✓ this results in about 8500 simulations
- ✓ 21 Mh cpu
- ✓ 24 Milions of files produced
- √ 54 TB of space
- ✓ using vectorization of 3D files –without that we would have 240 TB!

Some considerations ...

Model I/O

The run at BlueGeneQ facility of CINECA, where I/O is a bottleneck. Vectorization has been a crucial solution to:

- √ save disk space
- ✓ Save time both in simulation and in postproc
- ✓ Avoid zip/unzip (time consuming)

POSTPROC

Output must be reduced size in order

- ✓ to be stored and moved elsewhere
- ✓ to investigate easily on local servers the relationships parameter-variable
- 1 file for simulation (~10000 files)
- 1TB of space → that's enough for science

Simulation

Parallelism Level 1: Pure MPI Code, Mediterranean Sea divided in 128 tasks

Parallelism Level 2: Since the BlueGeneQ minimal requirement is 2048 cores, we used the SUBBLOCK task-farm technique. A job launch 16 different mpirun using a subset of cores

Postproc

With job of 2048 cores we process results of 165 simulations x 122 variables, i.e. we do 20000 times the same statistical work with the task-farm technique.

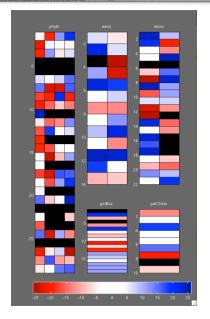
Solution adopted to remain on filesystem : 6D array

PROFILES [var, time, subbasin, distance from coast, z, kind of statistic]

No communication between ranks, every rank writes its file, then a serial collector writes the final file

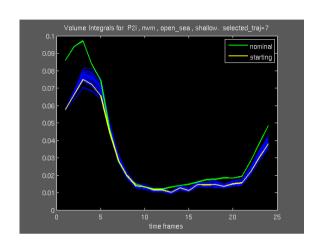


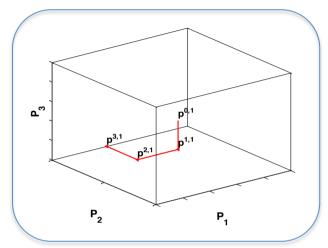
GSENSMED - Results

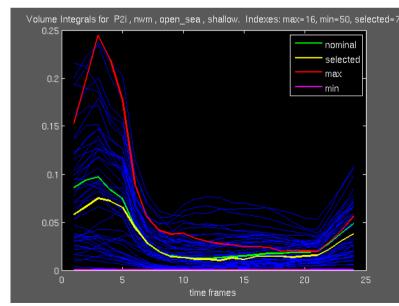


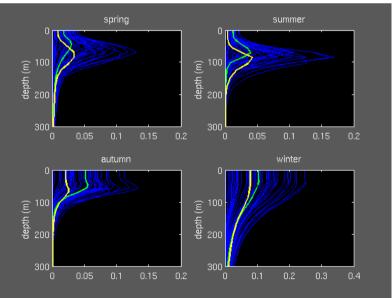
← Perturbation map of the namelist file







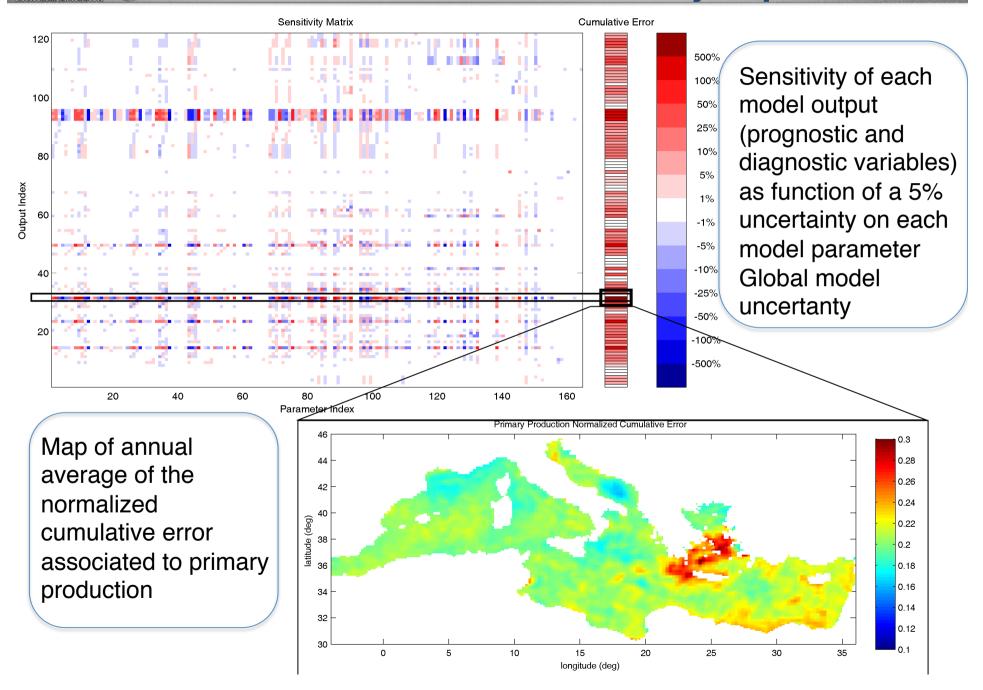




Results for the initial trajectories (51)



GSENSMED: Global sensitivity experiment





A 5% uncertainty in the value of each of the hundreds model parameters, with an estimate of the cumulative effect of such uncertainty.

Globally, the uncertainty is higher (up to 300%) in measures related to higher trophic level organisms, such as carnivorous mesozoopkanton, and lower for those related to planktonic primary producers

Several outputs present pretty large uncertainty, but –globally- the cumulative effect of potential uncertainty in hundreds of model parameters add up to an error which is smaller than 25% for the majority of the parameters.

In detail, the uncertainty associated to politically very relevant quantities, such as total primary productivity (a first proxy of sea fertility) or air-sea CO₂ fluxes (a first proxy of carbon sequestration) are about 5%.

The OFFLINE coupling approach

Advantages

- ✓ Reduction of the computational load especially needed in application with O(100) tracers to be advected (cell merge approach);
- √ capability to link with different OGCM (Operational or climate);
- ✓ the Offline approach simplify the task to carry out sensitivity analysis of complex 3D coupled models;
- ✓ Results are essentially equivalent to TM forced with higher resolution model (Levy et al., 2013).

Disadvantages

- ✓ formulation of the Transport Model must be sufficiently congruent with the formulation of the OGCM;
- ✓ Loose high frequency effects
- √ feedbacks of biology versus physics are not represented;
- ✓ disk space needed to store forcings (for 140 yrs~3.2Tb).



... COUPLING THE REACTOR IN A MODULAR WAY...



Shading effect of bio on radiation



A decrease in chl from 1.0 to 0.1 mg m⁻³ results in a 10 fold increase of net solar flux at 20 m

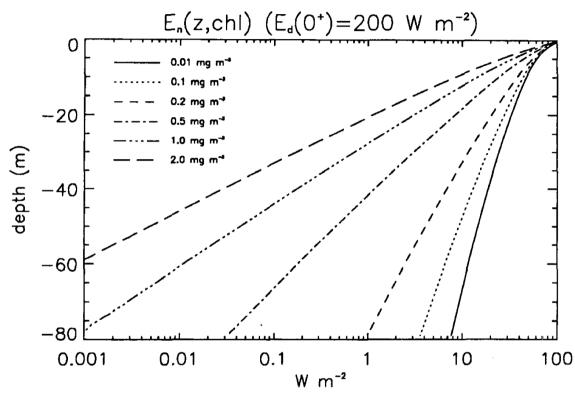


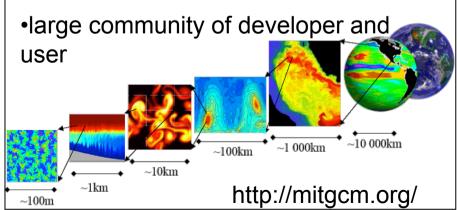
FIG. 3. Vertical profiles of net irradiance for an incident solar flux of 200 W m⁻² and chlorophyll a concentrations ranging from 0.01 to 2.0 mg m⁻³ after spectral integration from 300 to 750 nm.

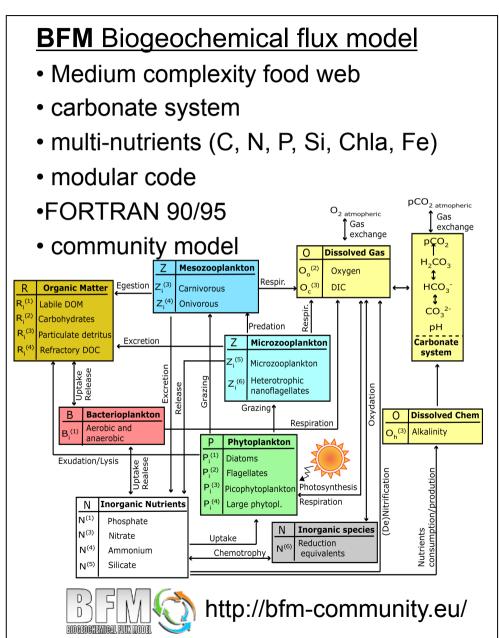
Ohlmann et al, 1996

COUPLED PHYSICAL-BIOGEOCH. MODEL

MITgcm: M.I.T. General Circulation Model

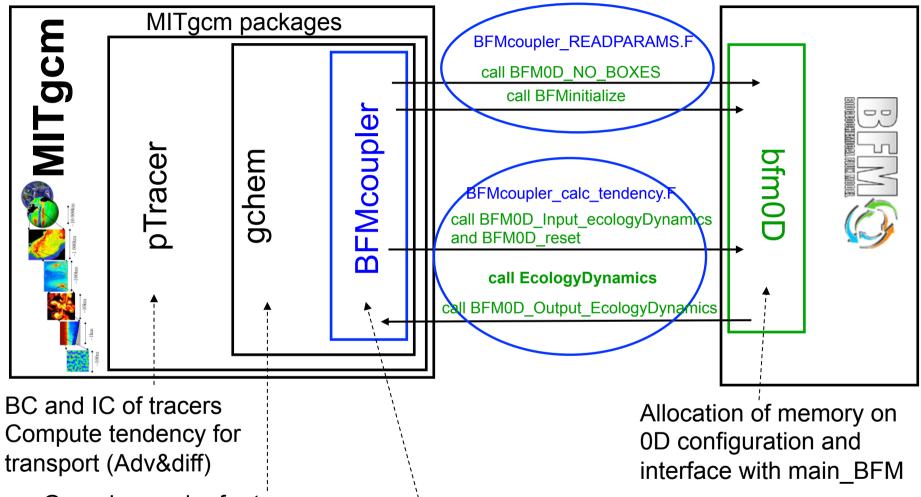
- •designed to study both atmospheric and oceanic phenomena, at various scales (from meter to global)
- includes non-hydrostatic capability
- adopts the KPP vertical turbulence parametrization
- adopts a finite volume technique
- developed to perform efficiently on a wide variety of computational platforms
- FORTRAN 77







Coupling MITgcm-BFM



Generic coupler for tracers call BFMcoupler routines (read params, external fields, initialization fields, and diagnostics)

Load and memory store of environmental local variables (P and N Deposition, atmpCO2, atmP, atmWind, surfPAR, PAR and light extinction factor) and biogeochemical tendency of tracers



... Some technical problems ... (to introduce the lab session)



- ✓ Is it possibile to couple a particular complex postproc procedure in the main model algorithm?
- ✓ Is it possible to threat the I/O procedures as a separate model to optmize the model performance?
- ✓ ... coupling a model with itself ...



Including the DA in the model framework



7 days of hindcast/analysis (using INGV physical forcing analysis and ICs via DA based on GOS-ISAC-CNR satellite chlorophyll)



10 days of forecast (using INGV physical forcing forecast)

Initial Conditions from previous run

3DVAR scheme using OC TAC data

Wednesday run

M TWTFSSMTWTFSSMTWTFS

TWTFSSMTWTFS

TWTFS

TWTF

TWTFS

TWTF

TWTFS

TWTFS

TWTFS

TWTF

T

Initial Conditions from previous run

Saturday run T F S S M T W T F S S M T W T F

Data Assimilation scheme (DA)

3D variational approach

 Cost function to obtain the innovation

Dobricic and Pinardi, 2008

$$J(\delta \mathbf{x}_k) = \delta \mathbf{x}_k^T \mathbf{B}_k^{-1} \delta \mathbf{x}_k + (\mathbf{d}_k - \mathbf{H}_k \delta \mathbf{x}_k)^T \mathbf{R}_k^{-1} (\mathbf{d}_k - \mathbf{H}_k \delta \mathbf{x}_k)$$

Model error covariance decomposition

$$\mathbf{B} = \mathbf{V}^T \mathbf{V} \qquad \delta \mathbf{x} = \mathbf{V} \mathbf{v}$$

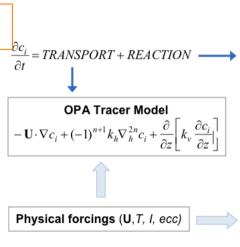
Solution in the control space v

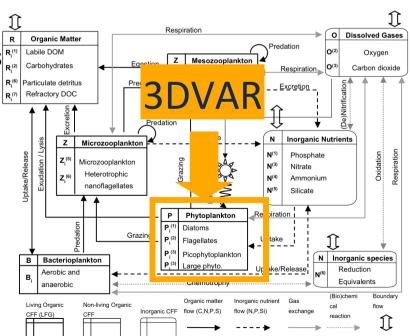
$$J(\mathbf{v}) = \mathbf{v}^{\mathrm{T}}\mathbf{v} + (\mathbf{d} - \mathbf{H}\mathbf{V}\mathbf{v})^{\mathrm{T}}\mathbf{R}^{-1}(\mathbf{d} - \mathbf{H}\mathbf{V}\mathbf{v})$$

V decomposed into a series of operators

$$\mathbf{V} = \mathbf{V}_{\mathbf{b}} \mathbf{V}_{\mathbf{h}} \mathbf{V}_{\mathbf{v}}$$

Teruzzi et al., 2014

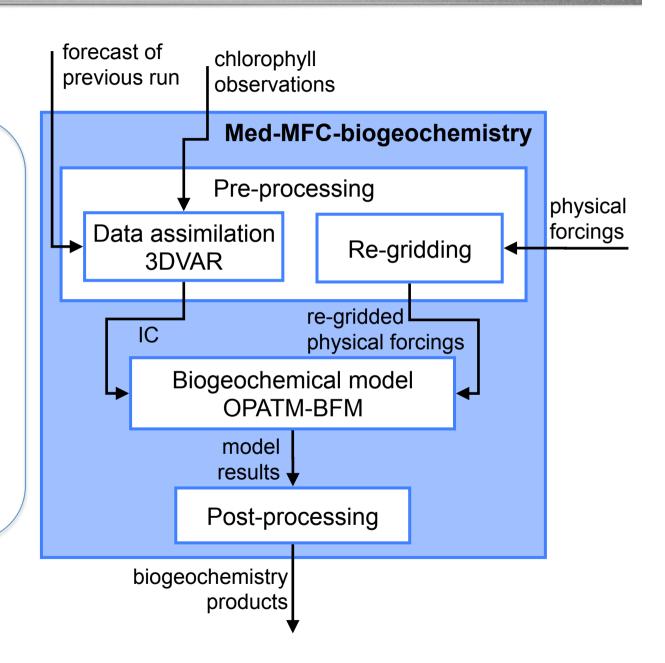






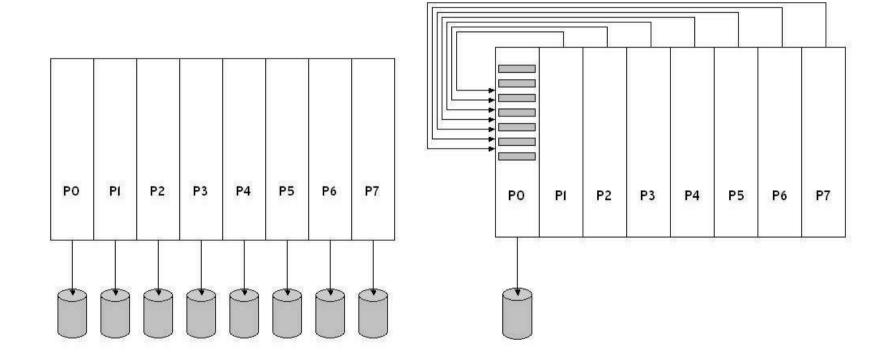
COUPLING THE DA WITH THE MODEL

- ✓ To perfom a reanalys long sequences of run followed by assimilation
- √To avoid reinitialization
 of the model coupling of
 the transport and DA
 model
- ✓ If both are parallel can they "coexist in a single MPI framework"?





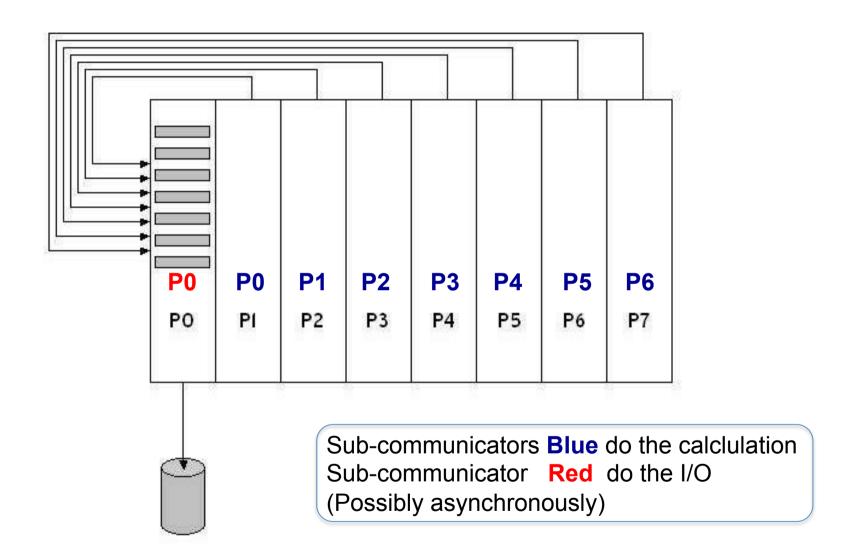
The problem of the I/O



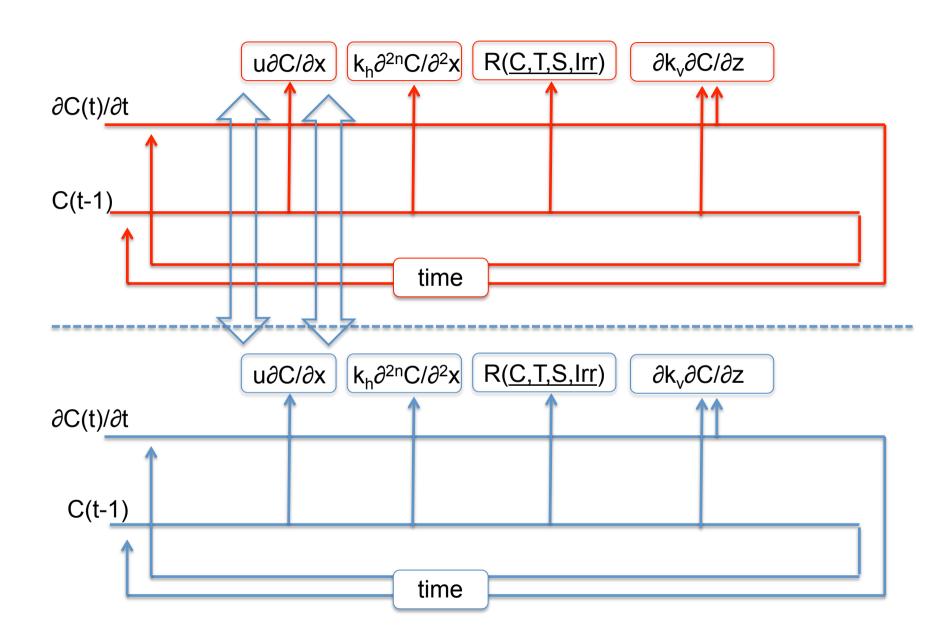
When P0 does I/O the others wait



Using different communicators









Thank you for the attention!