

The science of ocean predictions and downscaling to the coastal areas

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### Outline

- The science of ocean predictions: an historical viewpoint
- The Copernicus Marine Environment Service, analysis and forecasts at regional scales and forecast uncertainty estimation
- The downscaling conundrum: can we forecast better with limited area, nested models?

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Conclusions and Outlook

# The science of ocean predictions: historical viewpoint

- Bjerknes (1904, 1914) defined for the first time the 'rational method for weather predictions'
- Two conditions should be fulfilled in order to solve the prediction problem in atmosphere and oceans
  - I- Know the present state of the system as accurately as possible
  - II- Know the laws of physics that regulate the time evolution of the basic field state variables, i.e. have predictive models for atmosphere and oceans
- These concepts are at the basis of ocean prediction science also today

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#### The first ocean prediction: Harvard and Monterey 1983



Robinson (1983) defined the correct sampling scheme for Initialization (condition I) and selected the correct mesoscale model (condition II)



ROBINSON, CARTON, PINARDI AND MOOERS



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### The first ocean prediction: Harvard and Monterey 1983



Initial condition -



5516



5518

5512

5520



SEPTEMBER 1986



**Final forecast** 









# GOOS:operational oceanography starts in the 90s



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### **Dernicus** The European Earth Observation Programme The modelling component

A) hydrodynamics (1/16 x 1/16 x 72)

**B) Waves**  $(1/16 \times 1/16 \times 30)$ 

**C)** Pelagic biochemistry (1/16 x 1/16 x 72)





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#### Mediterranean subsystem: the quality component

RMS of Temp misfit at 8 m



2012

2013

2014

1.8

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2015

100

Number of measurements

Number of measurements

60

20



#### Vertical analysis error standard deviation for temperature and salinity TEMPERATURE SALINITY α α -30 -30 -60-60 -90 -90 -120-120 -150-150-180-180STD of model output -210 -210 STD of model output -240 -240 -270 -270 - 300 -300 SÉP MAY JÚN JÚL AÙG OCT MAY JÚN JÜL AUG SÉP DEC **FEB** NÖ DEC **FEB** D) D) 0.25 0.35 0.15 0.2 0.3 0.5 0.751.25 1.5 1.75 -30 -30 -60-60-90 -90 -120 -120-150 -150 -180 -180 -210-210Perturbed winds Perturbed winds -240-240 -270-270 200 JÜN AÙĠ SÉP OCT **FFB** SÉP MAY JUL Ν AÙG OCT MAY JUN JUL NÖ DEC FEB DEC 0.015 0.02 0.025 0.03 0.035 0.15 0.25 0.3 0.2

Errors in atmospheric forcing are projecting on the vertical structure of the temperature & salinity errors

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# Forecast Uncertainty estimation

- Shukla (2005): 'The largest obstacles in realizing the potential predictability of weather and climate are inaccurate models and insufficient observations, rather than an intrinsic limit of predictability'
- Uncertainty of ocean forecasts depends on:
  - Ocean Initial condition errors
  - Atmospheric forcing errors
  - Model errors (Physics, numerics)
- Hypothesis:
  - We use ensemble forecasting as a means to test ocean predictability issues
  - We concentrate on atmospheric wind forcing errors and how they affect the initial condition and forecast errors
  - We estimate realistic distribution of wind perturbations



#### Building the wind distributions using Bayesian Hierarchical Modelling (BHM-SVW)

Conceptual and implementation blocks:

Data Stage: 2 types of data, scatterometer winds and ECMWF analyses/ forecasts

Process model stage: Raylegh friction surface model translated into a stochastic finite difference equation





$$u = -\frac{f}{\rho_0 \left(f^2 + \gamma^2\right)} \frac{\partial p}{\partial y} - \frac{\gamma}{\rho_0 \left(f^2 + \gamma^2\right)} \frac{\partial p}{\partial x}$$

$$v = \frac{f}{\rho_0 \left(f^2 + \gamma^2\right)} \frac{\partial p}{\partial x} - \frac{\gamma}{\rho_0 \left(f^2 + \gamma^2\right)} \frac{\partial p}{\partial y}$$

$$U_t = \theta_{uy} D_y P_t + \theta_{ux} D_x P_t + \epsilon_u$$
$$V_t = \theta_{vx} D_x P_t + \theta_{vy} D_y P_t + \epsilon_v$$

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#### Posterior distributions of BHM-SVW (Milliff et al., 2011)







### The BHM-SVW Ocean Ensemble Forecast method (Pinardi et al., 2011)







at

he forecast spread



omv

ECMWF Ensemble Prediction System forcing is not effective to produce flow field changes at the mesoscales while BHM-SVW is

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#### forecasting and new applications





# The limited area forecasting conundrum

- The question is: can limited area ocean models increase coarse resolution forecast accuracy?
- Limited area forecasting requires to consider:
  - Coastlines at high resolution
  - High resolution bathymetry
  - Estuary forcings
  - Shelf break dynamics
  - The initialization problem
  - The lateral boundary condition problem
  - Surface atmospheric forcing of adequate resolution and extrapolated to minimize land contamination

that, very special problem

Need to





### Initialization problem: the spin up time

 Determination of spin-up time for the nested model (Simoncelli et al., DAO, 2013, De Dominicis et al., 2014)

From a 6 km model

Red-> nested 2 km model Blue-> nested 3 km model

Spin up time depends on the model domain and circulation regime but it is approx. **3-5 days** 



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#### Initialization problem: the interpolation from coarser resolution models

• Re-gridding of coarser fields in the finer grid (De Dominicis et al. OCDYN, 2013)



• coarse grid and velocity interpolated from coarse grid

Black: extrapolated from coarse grid with Viscous boundary layer assumption (AVERAGE OF NINE COARSE RESOLUTION GRID POINTS)

Variational methods can also be used to enforce no-slip boundary conditions at the wall after extrapolation

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# Lateral boundary condition problem

A For the tracers and total velocities at outflow/inflow:

$$\frac{\partial \theta}{\partial t} + u_{n \, coarse} \, \frac{\partial \theta}{\partial n} = \lambda (\theta - \theta_{coarse}) \quad ; \theta = (T, S, U_{total})$$

B For the barotropic component of velocity field, new GENERALIZED FLATHER BOUNDARY CONDITION has been developed (Oddo and Pinardi, 2008)

$$U_N^F = \frac{H_C + \eta_C}{H_F + \eta_F} U_N^C - \frac{C_N}{H_F + \eta_F} (\eta_C - \eta_F) \quad (!!!)$$

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C For different topography at the open boundaries INTERPOLATION

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**CONSTRAINT** (Pinardi et al., 2003)

#### Case 1: coastal forecasting with unstructured grid models (Federico et al., 2015)



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#### Case 2: relocatable structured and unstructured model (SURF, Trotta et al., 2015) Main characteristics: 1) Increase resolution only when it is needed and add physics, adapted to local conditions 2) Few hours deployment 3) Multiple nesting 4) Short term forecasting **First Parent domain** 6.5 km Operational model Second parent domain 2.2 km model Child domain 700 m model Department of Physics and Astronomy, University of Bologna Sept 29, 2015

#### Case 2: relocatable structured and unstructured model (SURF, Trotta et al., 2015)



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#### Case 2: **SURF** model error reduction



Green: nested model better in RMSE than the father model Red: the contrary



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#### **SHIP ROUTING**





### **Conclusions and Outlook**

- European Copernicus Marine Environment Service started operations and will continue thereafter
- Analysis errors and forecast uncertainty are being quantified. Ensemble methods with atmospheric forcing perturbations is a basic method for oceanography
- Open and free global and regional operational products make possible limited area, coastal short term forecasting
- Limited area forecasting IMPROVES forecast skill near coastal areas and in the open ocean with both structured and unstructured models
- When DSS are coupled to limited area models results show sensitivities

