

# Physics & Wildfires modeling



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NESDIS/OSI NOAA-15 AVHRR HRPT RGB=CH3,CH2,CH1 05/11/2000 01:30 UTC

# Wildfire: a complex multi-scale problem

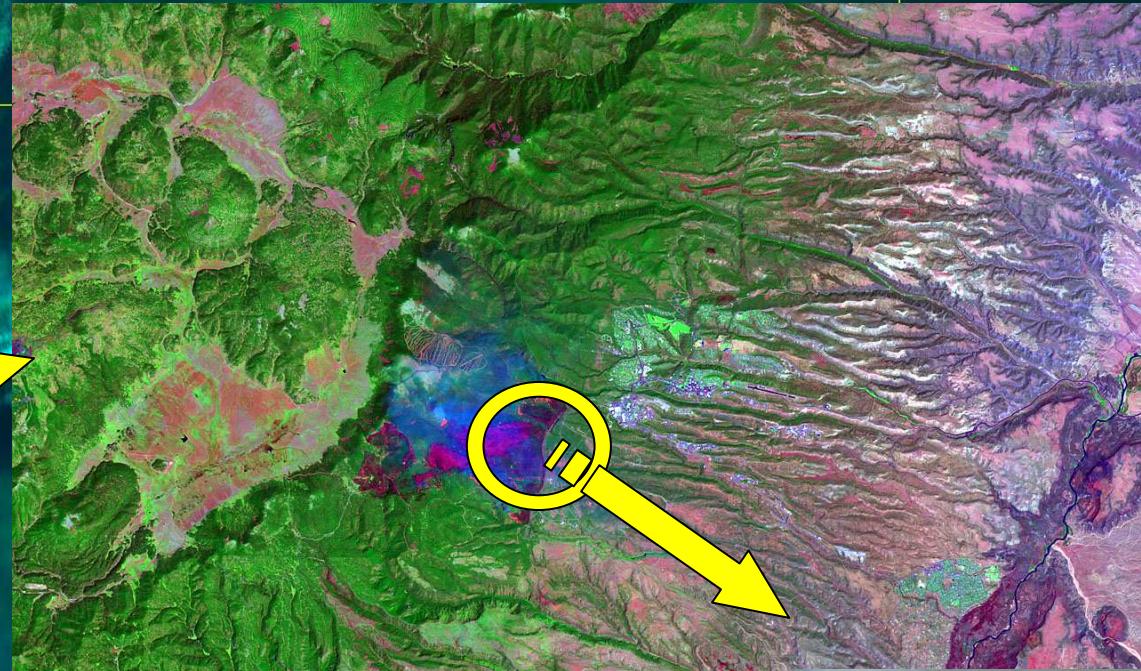
Colorado Kansas

(Los Alamos fire May 2000)

**140 km**

New Mexico

500 μm



# Wildfires: some physical characteristics (I)



## Flames:

- Turbulente,
- Very bright and sooty,
- Intense radiation,
- Soot production,

# Wildfires: some physical characteristics (II)

Rate of spread R (m/s): 0.1-1.5 m/s

Byram formula:

Fire line intensity I (kW/m):  $I = M_{burn} \times \Delta H$

USA, Australia: 100 000 kW/m

Europe: 10 000 kW/m

$M_{burn}$ : Solid fuel consumption rate (kg/m/s)

$\eta$ : Combustion efficiency

$M_{fuel}$ : Dry solid fuel load (kg/m<sup>2</sup>)

$\Delta H$ : Heat of combustion (~18 700 kJ/kg)

$H_f$ : Flame height (m)

$$I = \eta M_{fuel} \times \Delta H \times R \sim 300 \times H_f^2$$

# Fire intensity scales...



Efficiency limit: 2000 kW/m (terrestrial), 7000 kW/m (aerial)

# **Black Saturday (07/02/2009)**

## **Victoria district (Australia)**

**Fire storm (Kinglake)**

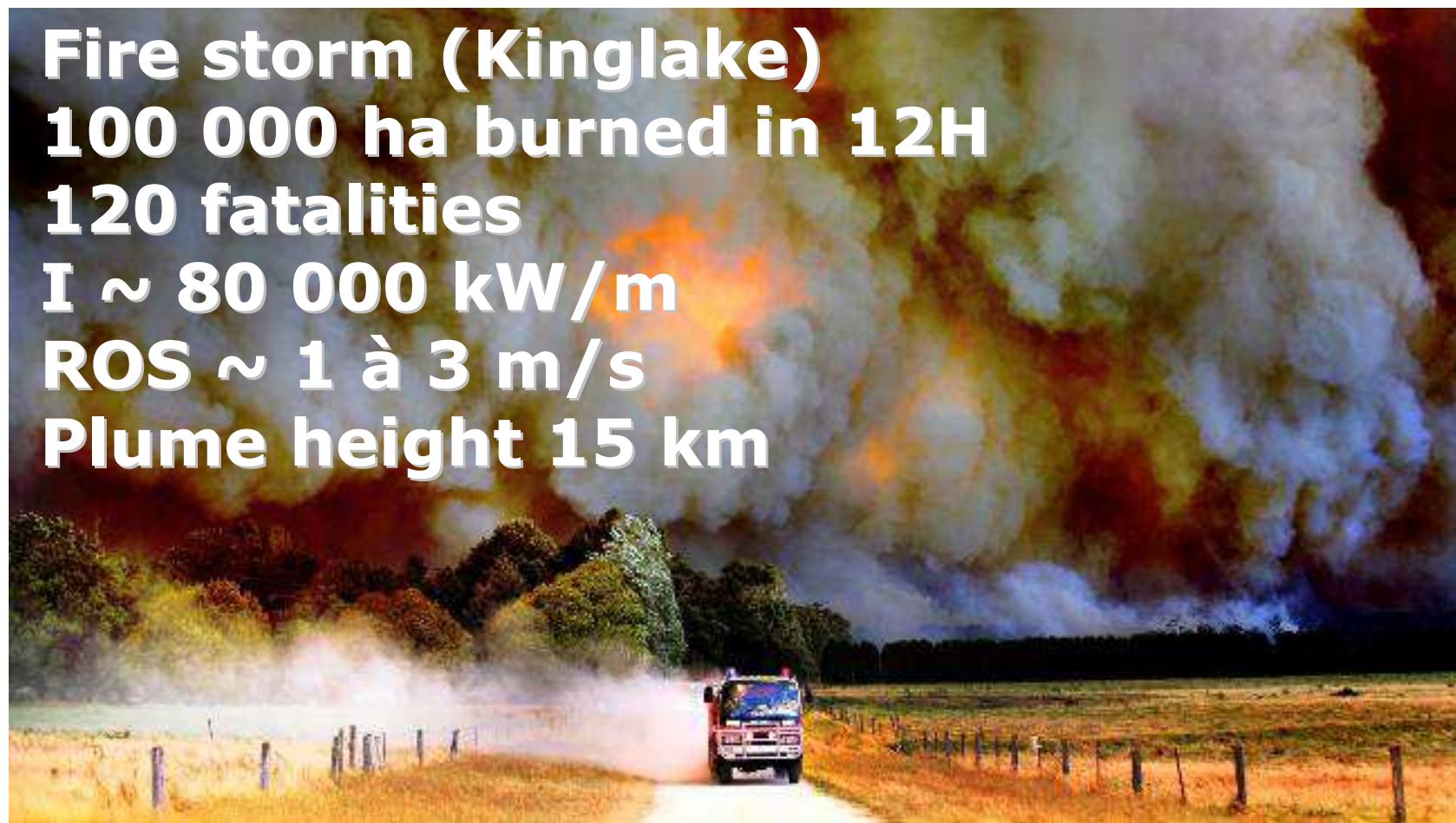
**100 000 ha burned in 12H**

**120 fatalities**

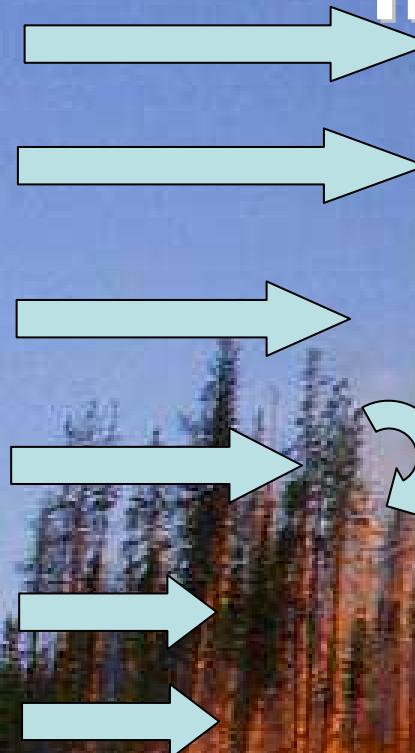
**$I \sim 80\ 000\ kW/m$**

**$ROS \sim 1 \text{ à } 3\ m/s$**

**Plume height 15 km**



# Wildfires modelling: a complex multi-scale problem



Soot + hot gases  
Radiation + Convection

Fire front (combustion, turbulence):  
Flame thickness:  $d_F \sim 500 \mu\text{m}$

Fuel: drying,  
pyrolysis,  
combustion  
 $H_{Fuel}$

ABL (turbulence):  
Large scale:  $L_t \sim H_{Fuel}$   
Micro scale:  $\eta \sim 100 - 500 \mu\text{m}$

Radiation heat transfer  
(length of extinction):  
 $L_R \sim 2xH_{Fuel} / LAI$   
0.1 – 5 m

# An other mecanism of fire propagation : firebrands



Distance travelled by a firebrand > 2400 m (source: SALTUS) !

# Factors affecting the behaviour of forest fires: the fire triangle



## Topography

- Flat or slopes
- Aspect

## Weather

- Wind
- Temperature
- Relative Humidity
- Precipitation

## Fuel

- Fine or Heavy
- Arrangement & continuity
- Fuel Moisture

# Atmosphere/wildfire interaction

Unstable



Neutral or stable



$$\frac{dT}{dz} < DATG$$

$$\frac{dT}{dz} \geq DATG$$

**DATG: dry adiabatic temperature gradient**

$$DATG = -\frac{g}{c_p} \approx -10 K / km$$

## Wildfires modeling

(Weber 1991, Sullivan 2009 ...)

- Statistical models (empirical),

→ Mc Arthur (1966)  $R = f(f_i)$

- Semi-empirical models,

→ Rothermel (1972)  $R = \xi I_r / \rho \Delta h_i$

- Physical models (radiative, full physics),

→ Albini (1985), De Mestre (1989),  
Balbi, Santoni & al (1998), Siméoni & al (2001)

$$R = q / \rho |dh/dx|$$

→ Grishin (1985), Larini, Morvan, Porterie & al (1996)  
Clark & al (1996), Lin & al (1997),  
Sero-Guillaume & al (2002), Rehm & al (2003), Mandel & al  
(2004), Mell & al (2005), Mahalingam (2008), Filippi & al  
(2009) ...

# Semi-empirical Rothermel's model de (BEHAVE, FARSITE)



Fuel:



Excelsior



Pine needles

$$R = \xi I_r / (\rho_s \alpha_s \Delta h_i), \quad \xi = (192 + 7.894 \sigma_s)^{-1} \exp[(0.792 + 3.760 \sigma_s^{1/2}) (\alpha_s - 0.1)]$$

$\rho_s \alpha_s$  : Fuel density and fuel volume fraction

$I_r$  : Heat of combustion

$\Delta h_i = C_p [T_i - T_a]$  Enthalpy of ignition

$\sigma_s$  : Surface area / Volume ratio of solid fuel particles

# Semi-empirical Rothermel's model de (BEHAVE, FARSITE)



$$R = R_0 [1 + \phi_w + \phi_s]$$

$$\phi_w = C_w U_w^{Bw} [\alpha_s / \alpha_s^0]^{-Ew} \text{ (wind effect)}$$

$$\phi_s = 5.275 \alpha_s^{-0.3} \operatorname{tg}^2(\varphi) \text{ (slope effect)}$$

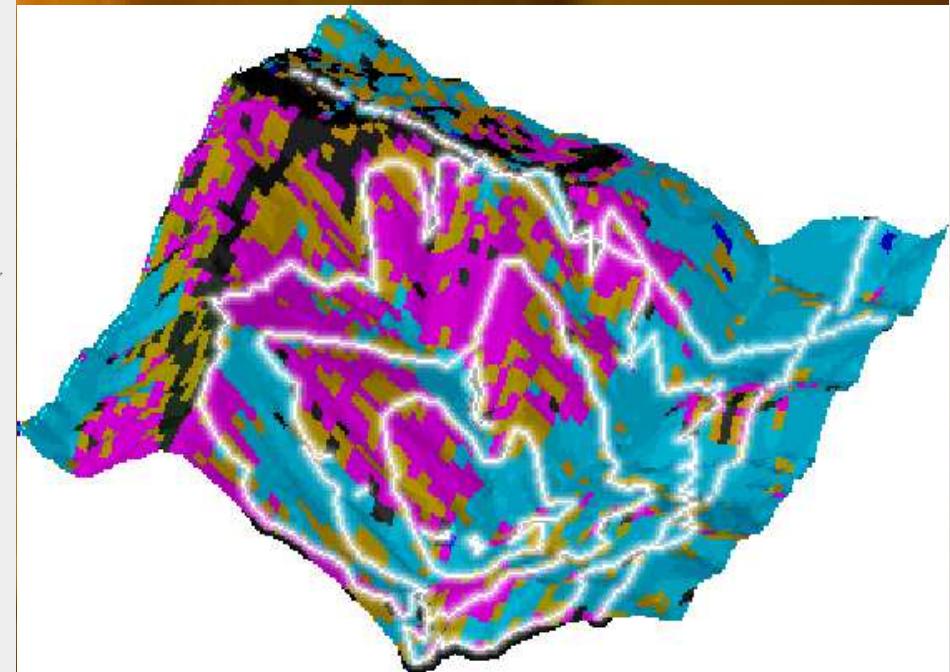
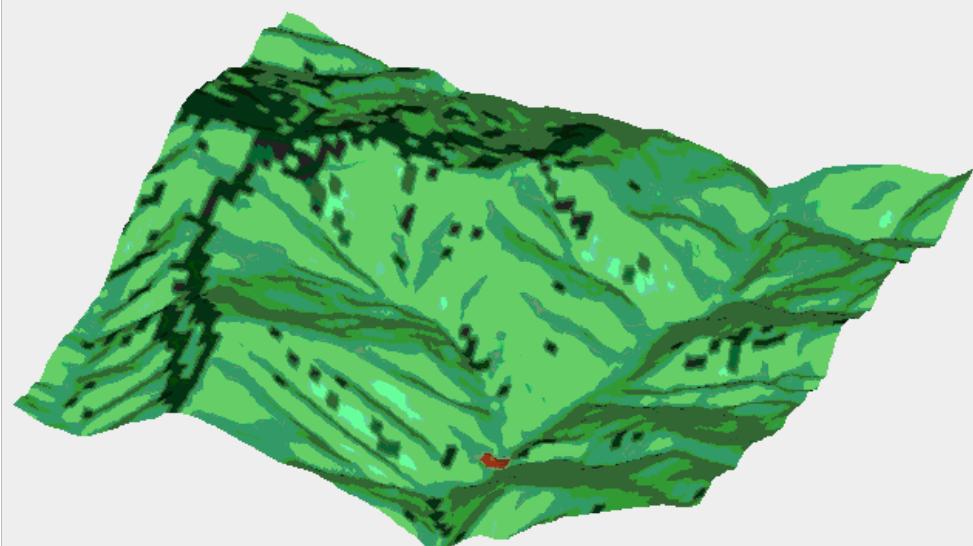
$$C_w, B_w, E_w = f(\sigma_s), B_w = 0.0132 \sigma_s^{0.54}$$

$$B_w = 0.4 (\sigma_s = 666 \text{ m}^{-1})$$

$$B_w = 1.6 (\sigma_s = 7596 \text{ m}^{-1})$$

# Malibu fire (22/10/1996)

AVIRIS Derived Fuel



HIGRAD-FIRETEC

FARSITE

Hanson & al Environ. Sci. Policy 2000

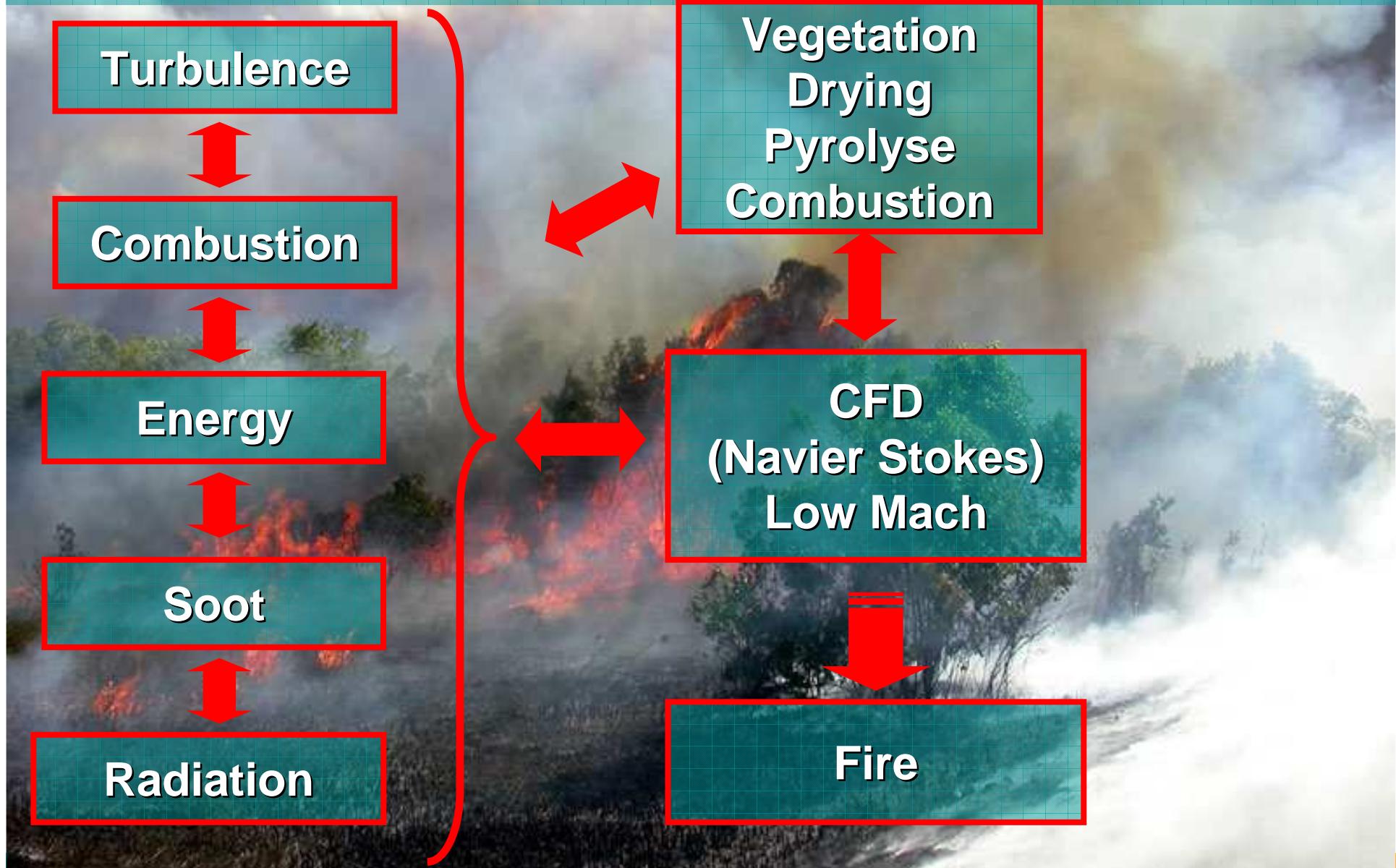
# Malibu fire: semi-empirical vs physical models

Time observed to burn 50 ha (real time = 10 minutes)  
(Hanson & al 2000)

... effect upon the wind	Slope	Fire	Time
Farsite (Behave)			180 min.
Behave+Higrad	X		20 min.
Firetec + Higrad	X	X	10 min.

Hanson & al Environ. Sci. Policy 2000

# Simulating wildfire behaviour using a physical model : multiphase approach



# ABL/canopy interaction



## Leaf Area Density

$$a_L = LAD = \frac{\alpha_s \sigma_s}{2}$$

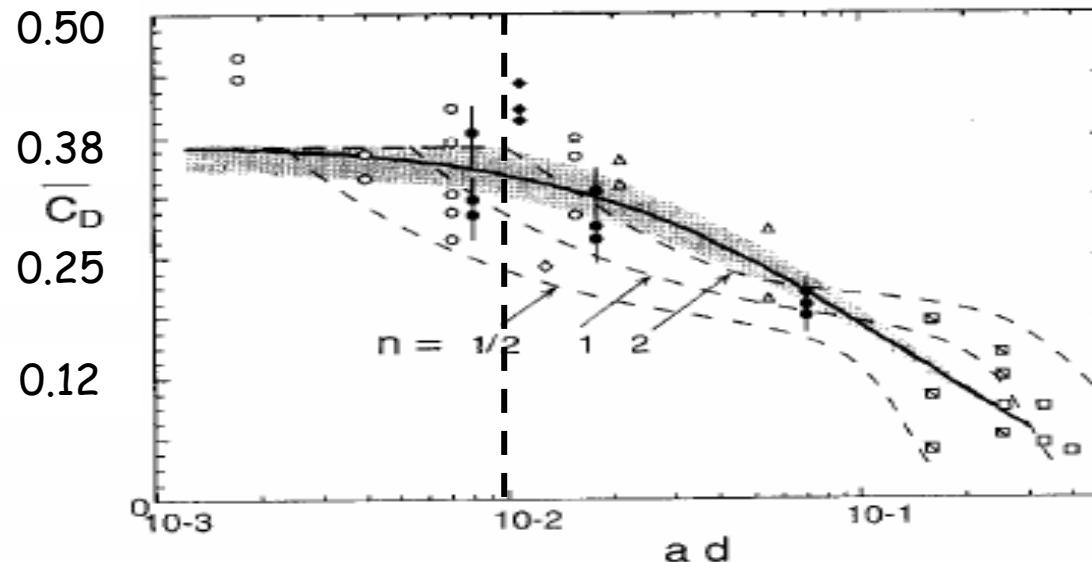
$$\frac{\partial}{\partial t}(\rho \langle u_i \rangle) + \frac{\partial}{\partial x_j}(\rho \langle u_i \rangle \langle u_j \rangle) = \frac{\partial \langle \sigma_{ij} \rangle}{\partial x_i} + \frac{\partial \tau_{ij}}{\partial x_i} - a_L C_D \rho |\langle u \rangle| \langle u_i \rangle$$

$$\frac{\partial \langle u' w' \rangle}{\partial z} = -C_D \times LAD \times \langle u \rangle^2 \Rightarrow C_D = 0.1 - 0.4$$

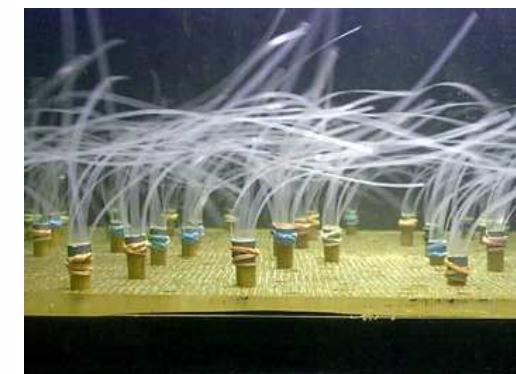
$$\frac{D\rho K}{Dt} = \frac{\partial}{\partial x_j} \left[ \frac{\mu_{eff}}{\sigma_k} \frac{\partial K}{\partial x_j} \right] + \rho P - \rho \varepsilon + a_L C_D \rho \left[ \langle U \rangle^3 - 4 \langle U \rangle K \right]$$

# Drag coefficient in a sea grass

( $C_D$  défini à partir de la LAD (Leaf Area Density))



$$ad = \frac{4\alpha_s}{\pi}$$



- If  $ad < 0.01$   $\langle C_D \rangle = C_D(R_e)$  (~ single particle)
- If  $ad > 0.01$   $\langle C_D \rangle = f(ad)$  (wake interaction)
- Typical values:  $ad \sim \alpha_s \sim 10^{-3} - 10^{-2}$

(Water Resources Research Vol.35(2) pp.479-489 (1999), H.M. Nepf)

# Energy balance in the solid phase

$$\alpha_{s,p} \rho_{s,p} C_{s,p} \frac{dT_{s,p}}{dt} = h_{conv} \underbrace{\alpha_{s,p} \sigma_{s,p} (T - T_{s,p})}_{\text{Convection}} + \underbrace{\frac{\alpha_{s,p} \sigma_{s,p}}{4} (J - 4\sigma T_{s,p}^4)}_{\text{Radiation}} - \sum_{\alpha} M_{s,\alpha} h_{s,\alpha}$$



**Convection**  
 $\sim$  linear



**Radiation**  
non-linear

# Turbulence/Radiation Interaction

## Optically thin fluctuation approximation (OPFA)

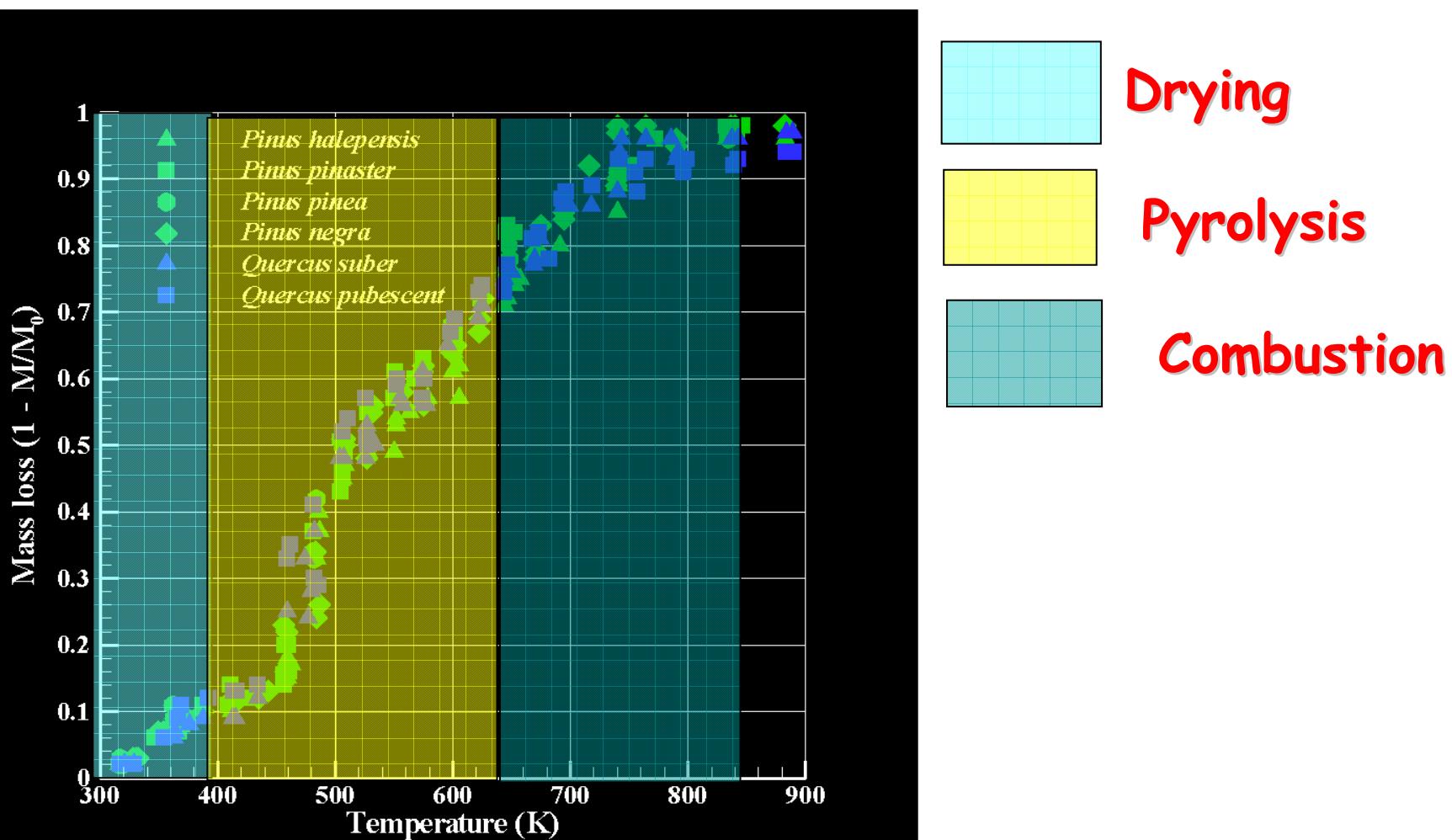
$$\frac{d\alpha_G \bar{I}}{ds} = \alpha_G \left( \frac{\sigma \bar{\sigma}_a T^4}{\pi} - \bar{\sigma}_a \bar{I} \right) + \frac{\sigma_s \alpha_s}{4} \left( \frac{\sigma T_s^4}{\pi} - \bar{I} \right)$$

$$\bar{\sigma}_a T^4 \approx \bar{\sigma}_a \bar{T}^4 \left[ 1 + 6 \frac{\bar{T}'^2}{\bar{T}^2} + 4 \frac{\bar{\sigma}_a' \bar{T}'}{\bar{\sigma}_a \bar{T}} \right] = \bar{\sigma}_a \bar{T}^4 \left[ 1 + 6 \frac{\bar{T}'^2}{\bar{T}^2} + 4 \frac{\bar{T}'^2}{\bar{\sigma}_a \bar{T}} \frac{\partial \sigma_a}{\partial T} \right]$$

$$\frac{\partial \sigma_a}{\partial T} = 1862 \times \alpha_{soot} \quad \bar{T}'^2 = \theta$$

$$\frac{D\bar{\rho}\theta}{Dt} = \frac{\partial}{\partial x_j} \left( \frac{\mu_{eff}}{\text{Pr}_T} \frac{\partial \theta}{\partial x_j} \right) + 2P_\theta - 2\varepsilon_\theta \quad P_\theta = \frac{\mu_T}{\text{Pr}_T} \left( \frac{\partial \bar{T}}{\partial x_j} \right)^2 \quad \varepsilon_\theta = \rho \frac{\theta}{2 \times R} \times \frac{\varepsilon}{K}$$

# Thermal analysis of Mediterranean vegetation samples (INRA-Avignon)



# Mass balance in the solid phase

## Water & dry solid fuel

$$\frac{d(\alpha_{s,p} \rho_{s,p} Y_{s,p}^{\text{H}_2\text{O}})}{dt} = -\omega_{vap}^s$$

$$\frac{d(\alpha_{s,p} \rho_{s,p} Y_{s,p}^i)}{dt} = -\omega_{pyr}^s$$

## Charcoal

$$\frac{d(\alpha_{s,p} \rho_{s,p} Y_{s,p}^{\text{char}})}{dt} = (v_{char} - v_{soot}) \omega_{pyr}^s - \left( \frac{v_{ash}}{v_{char}} + 1 \right) \omega_{char}^s$$

# Mass balance in the solid phase

## Global mass balance

$$\frac{d(\alpha_{s,p} \rho_{s,p})}{dt} = -\sum_{\alpha} M_{s,p,\alpha}$$

## Volume balance

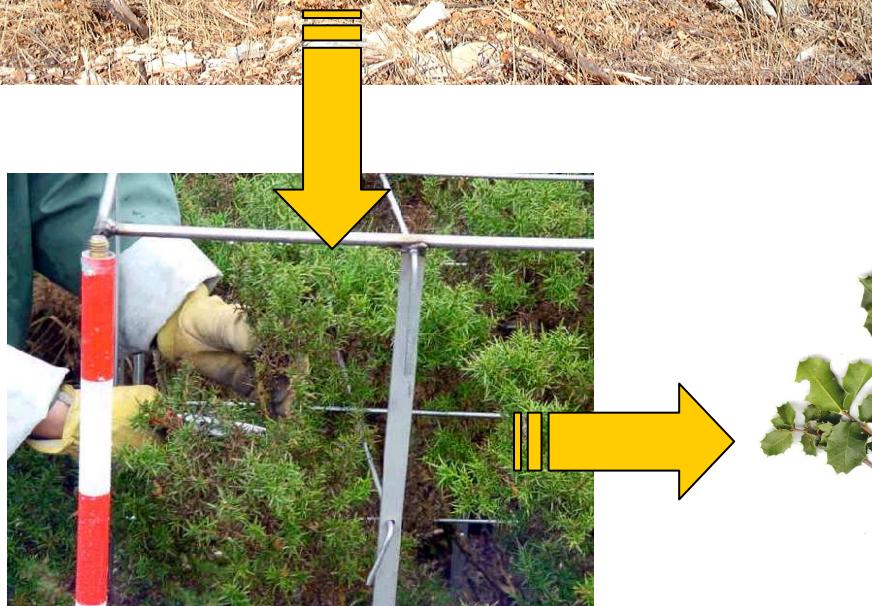
$$\frac{d\alpha_{s,p}}{dt} = -\frac{1}{\rho_{s,p}} \omega_{char}^s$$

# Physical description of the vegetation layer

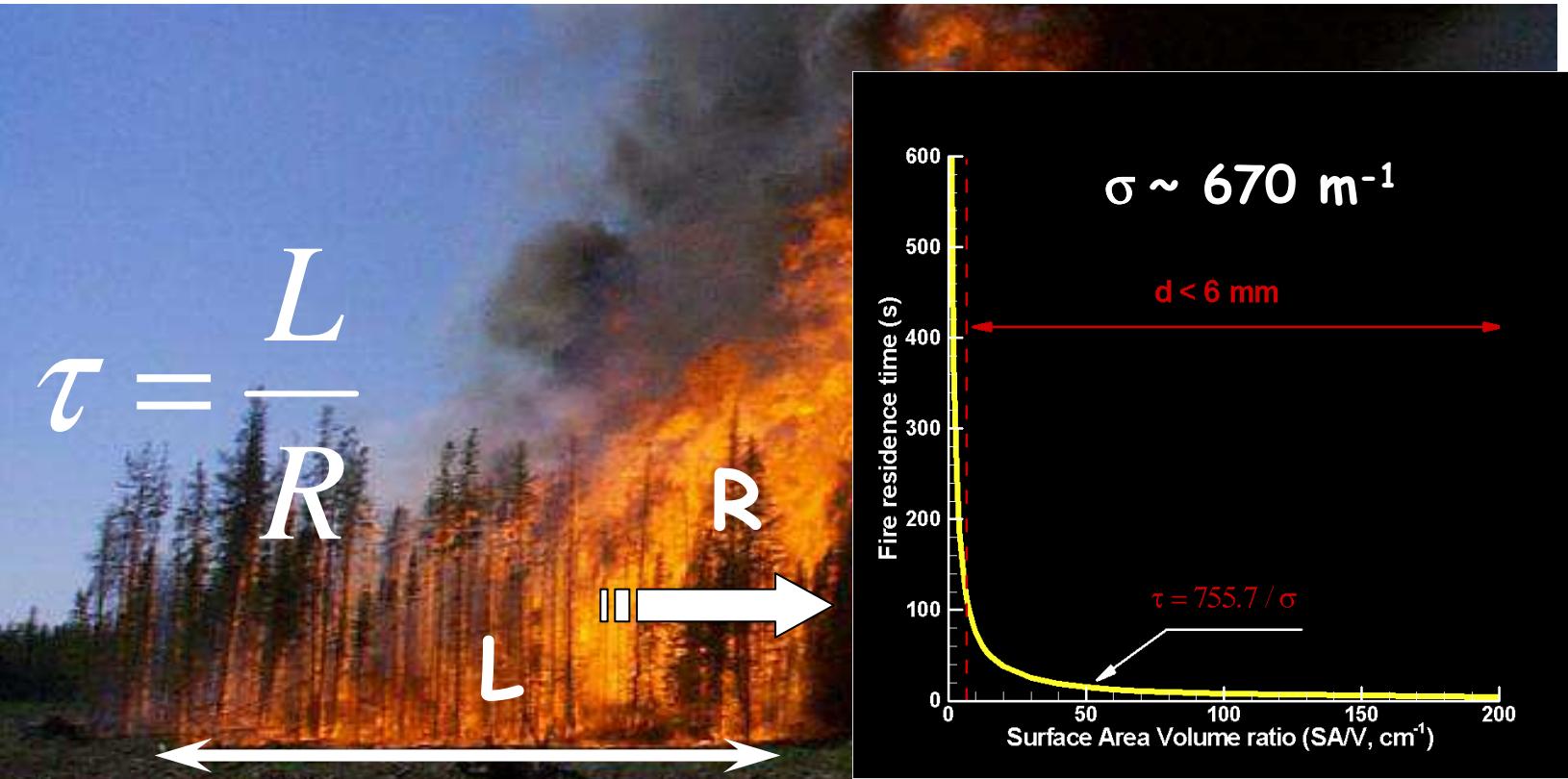


## Physical properties

- Density
- Volume fraction
- Surface Area/Volume (SA/V)
- Fuel moisture content (FMC)



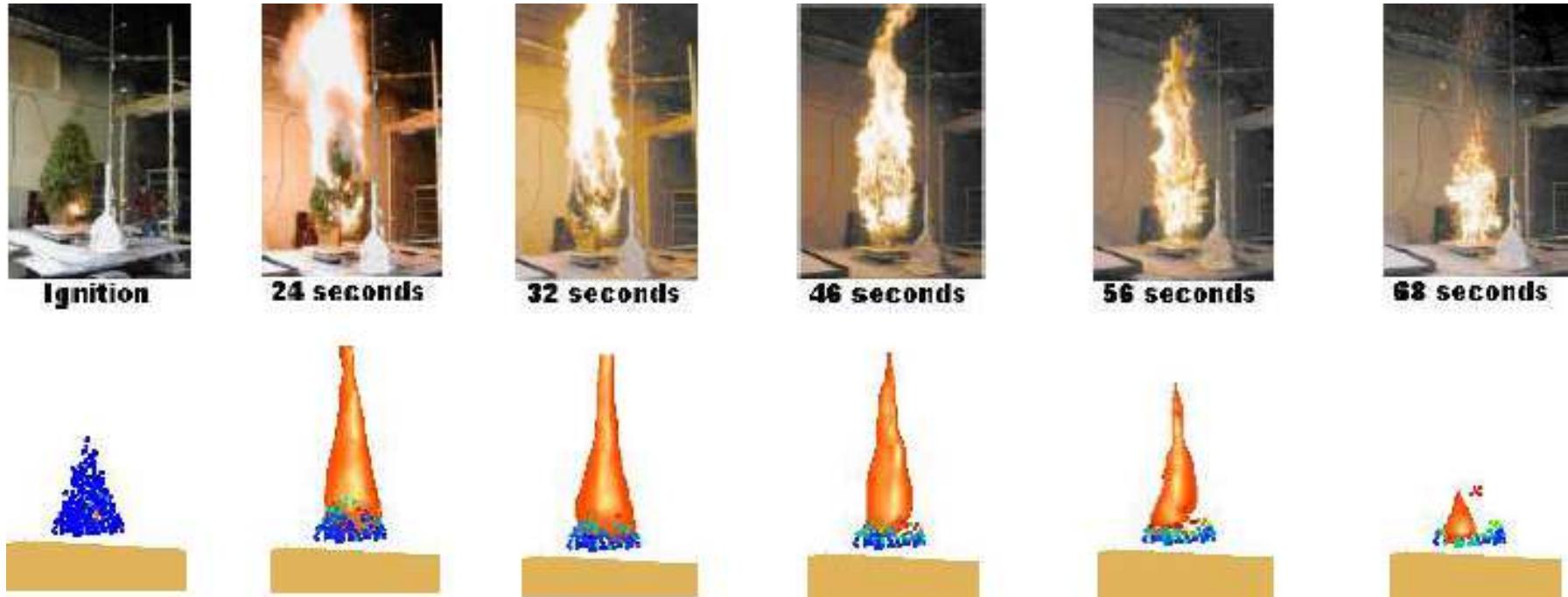
# Fire residence time



$\sigma (\text{m}^{-1})$	600	2000	5000	10000	20000
$\tau (\text{s})$	125	37	15	7	3

# New tools for simulating wildfires

## WFDS: Wildland Fire Dynamic Simulator

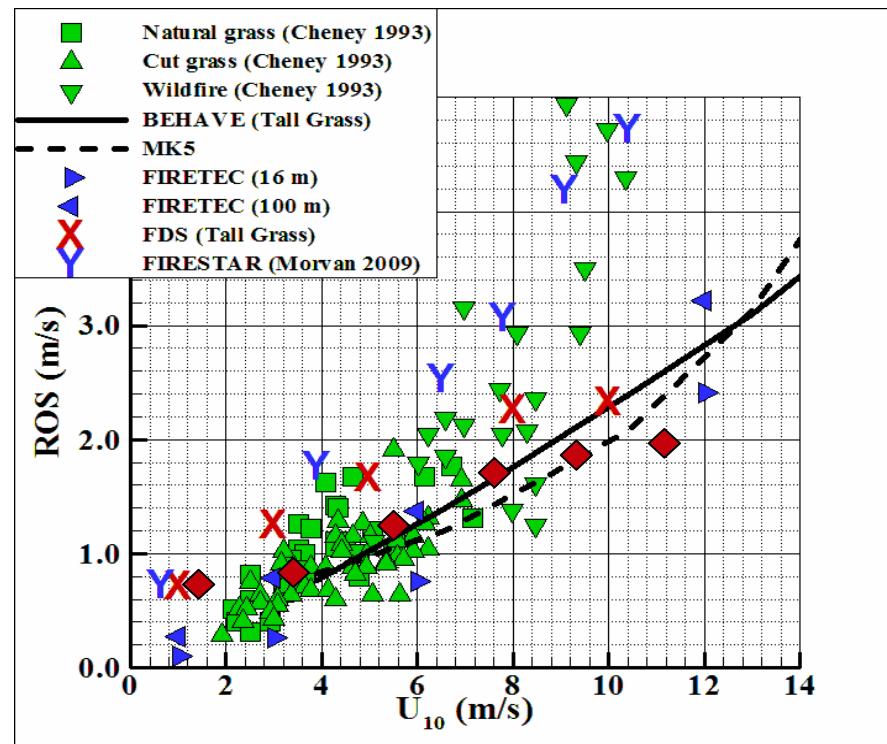


Mell & al Combust. Flame 2009  
US Forest Services Pacific Wildland Fire Sciences Lab. (Seattle)  
NIST National Fire Research Lab. (Gaithersburg)

# Large scale experimental fires (grassland, CSIRO, Australia)

**Plot: 20 m x 50 m (+ safety band)**

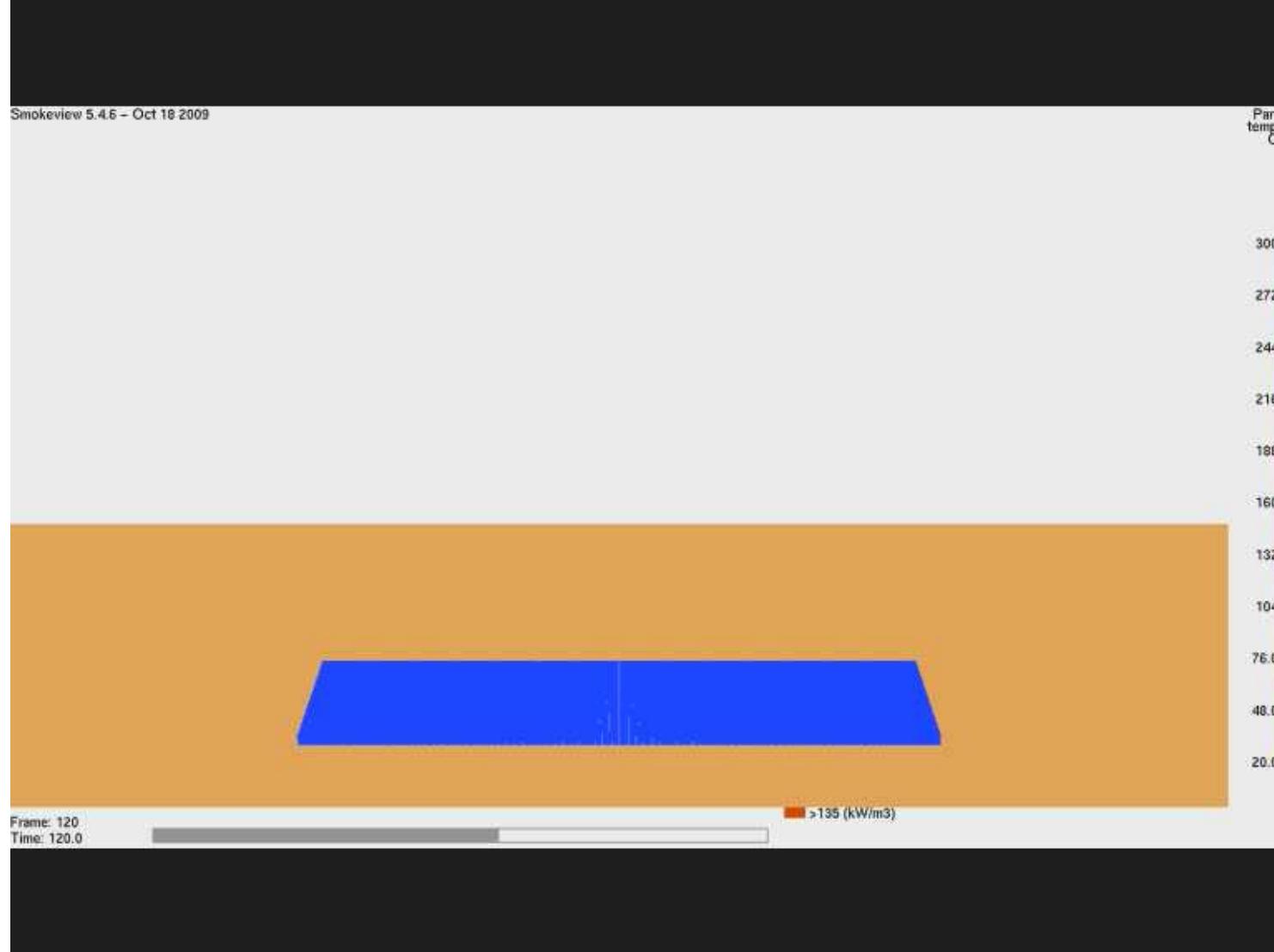
**Mesh > 20 Millions cells**



**Cheney & al Int. J. Wildland Fire 1998**  
**Morvan & al Fire Safety Journal 2009**

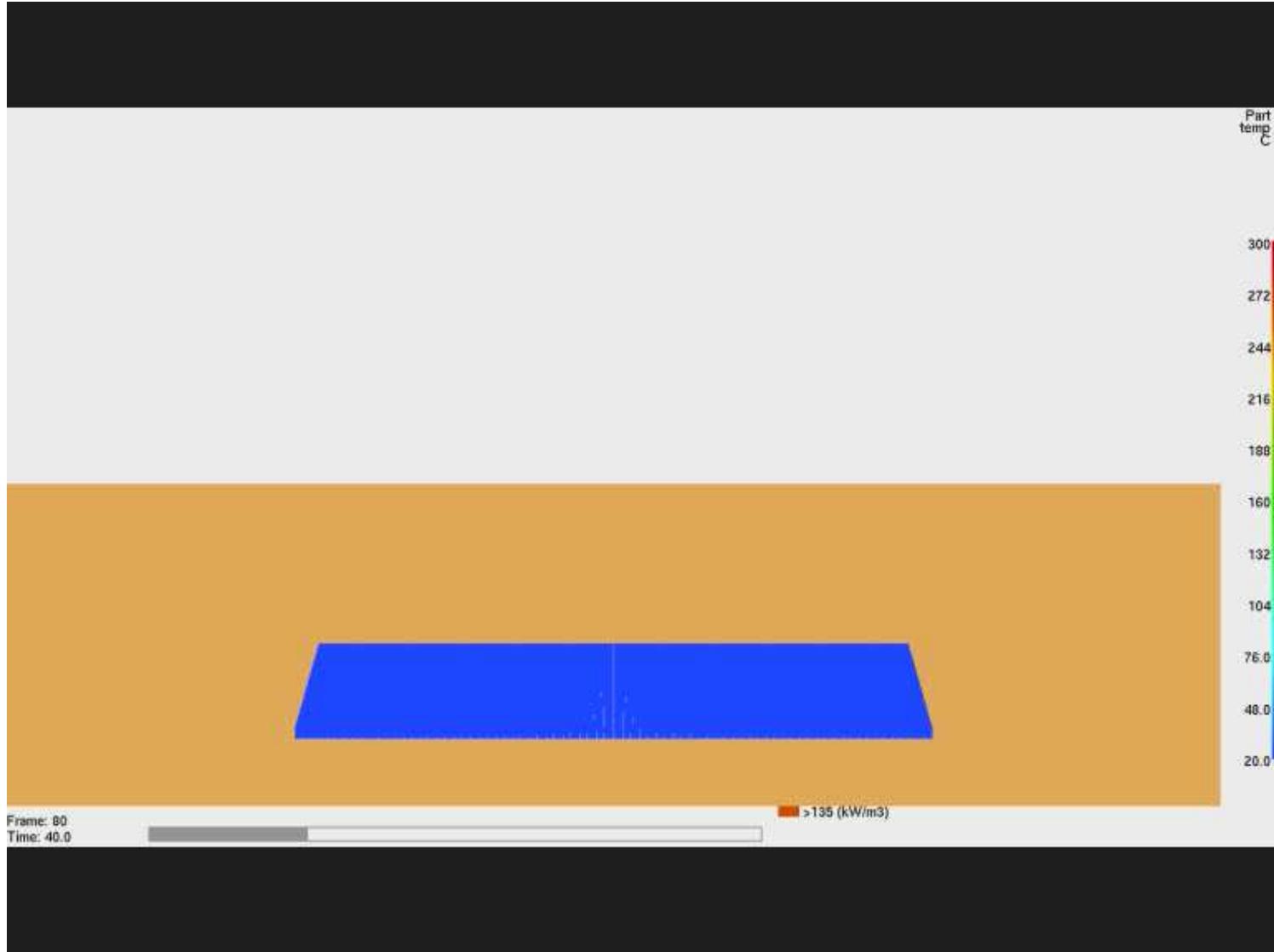
# Fire in grassland: 3D simulation (WFDS)

Plot: 50m × 20m  $U_{10} = 1 \text{ m/s}$



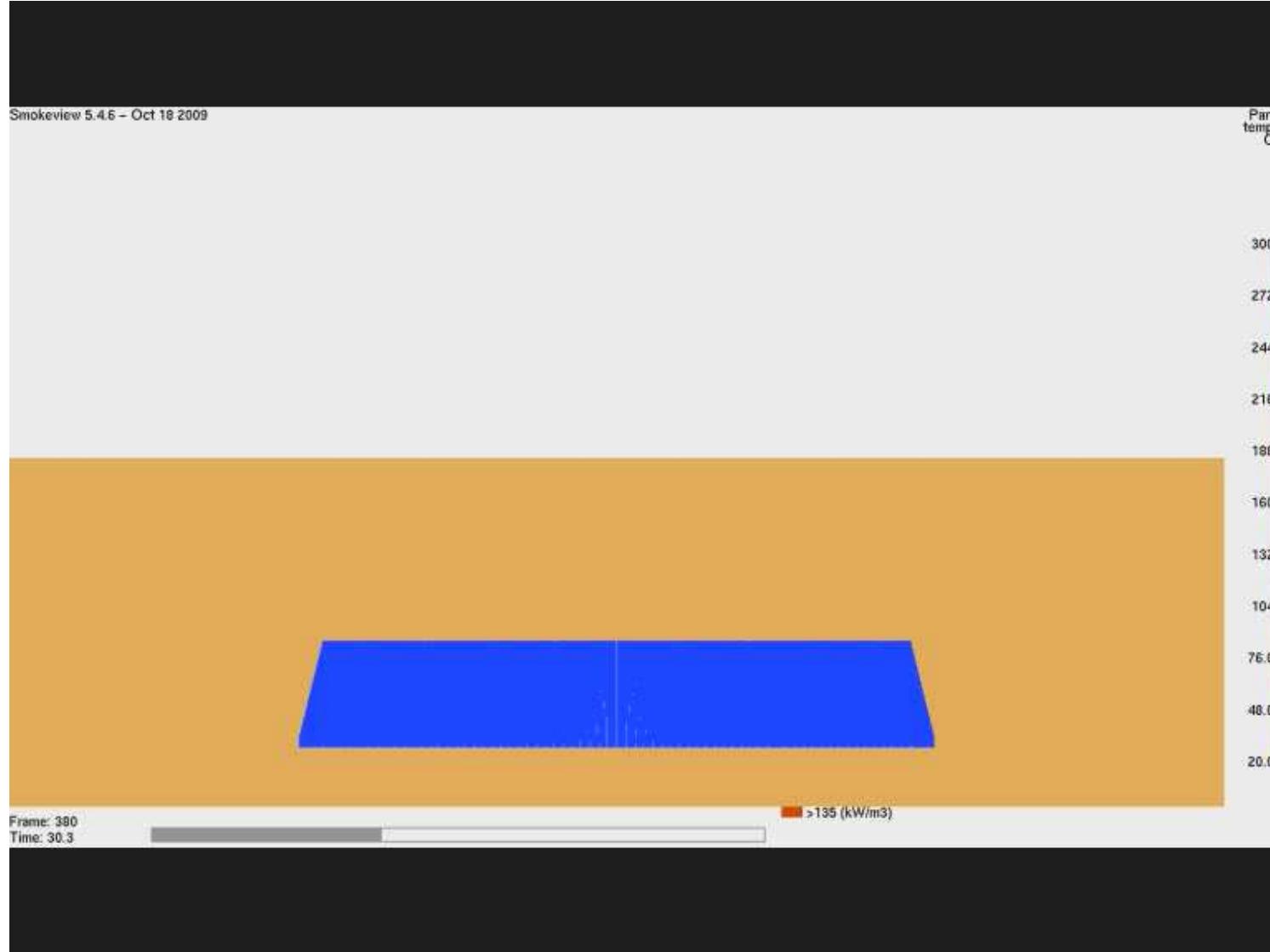
# Fire in grassland: 3D simulation (WFDS)

Plot: 50m × 20m  $U_{10} = 3 \text{ m/s}$



# Fire in grassland: 3D simulation (WFDS)

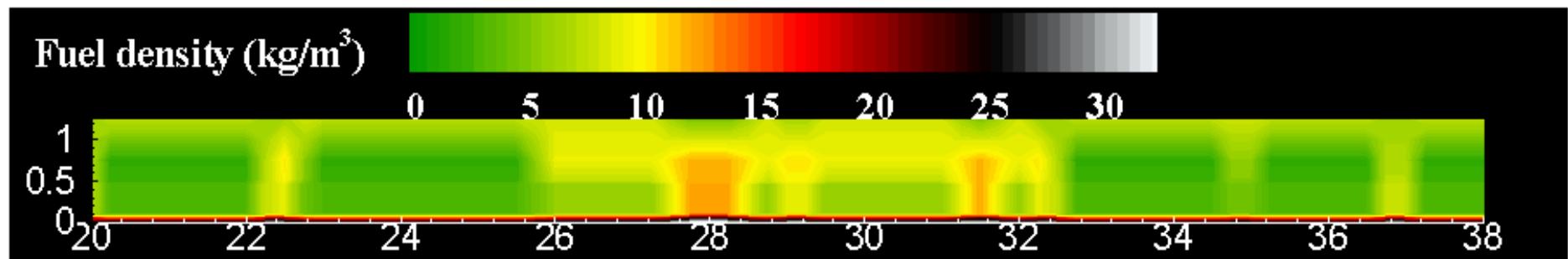
Plot: 50m x 20m  $U_{10} = 10 \text{ m/s}$



# Experimental fire in shrubland (EU Firestar project, Galicia-Spain)



# Experimental fire in shrubland (EU Firestar project, Galicia-Spain)

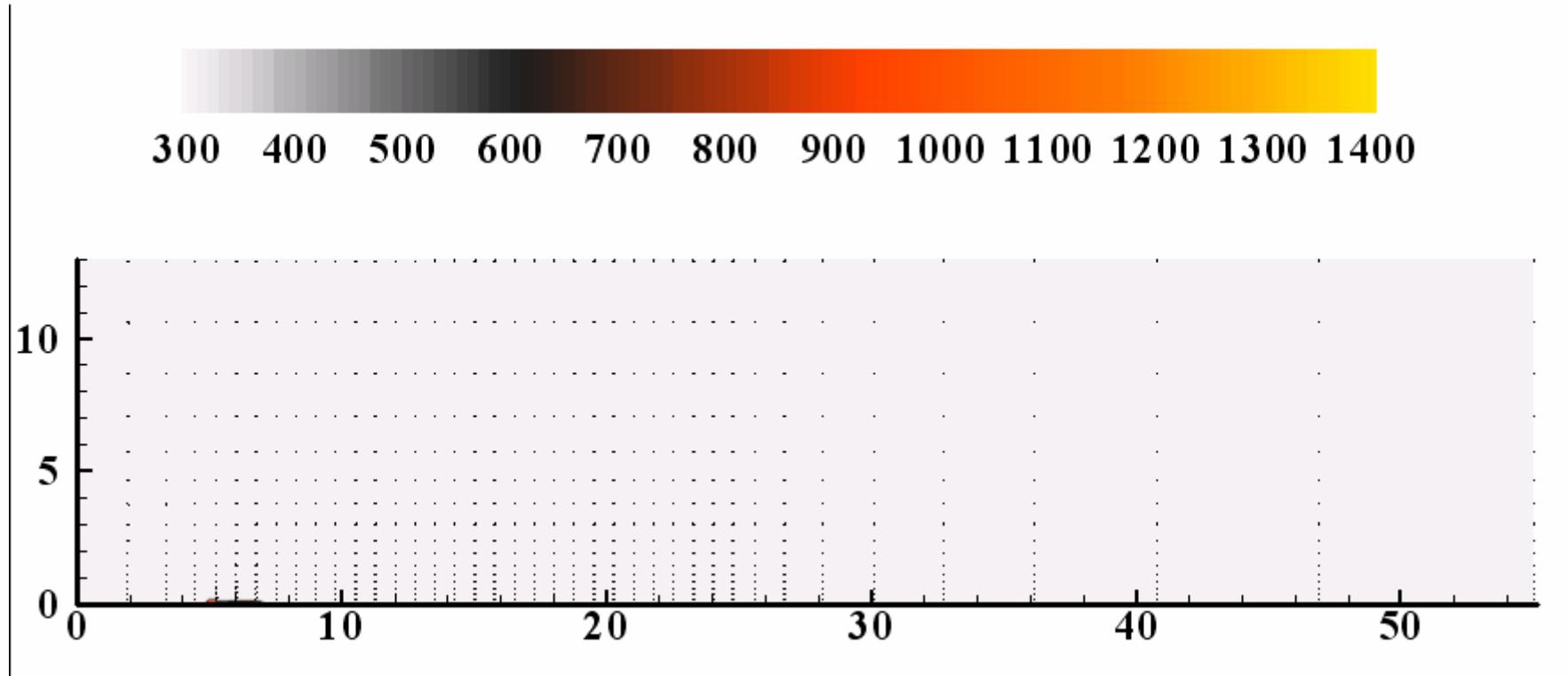


- Fuel: *Ulex (Europaeus, Minor)*,
- Fuel families=14
- FMC:  
**108-150 % (vivant), 10-32% (mort)**
- Fuel depth = 1.25 m,
- Wind : 5.7 m/s ( $z=10$  m),
- Slope : 5°



# Experimental fire in shrubland *(EU Firestar project, Galicia-Spain)*

Grassland fire | 30 Nov 2005 | FIRESTAR



- Experiment:  $ROS = 0.273 \text{ m/s}$
- Simulation :  $ROS = 0.248 \text{ m/s}$

# Wildfire propagation regimes Convective Byram number

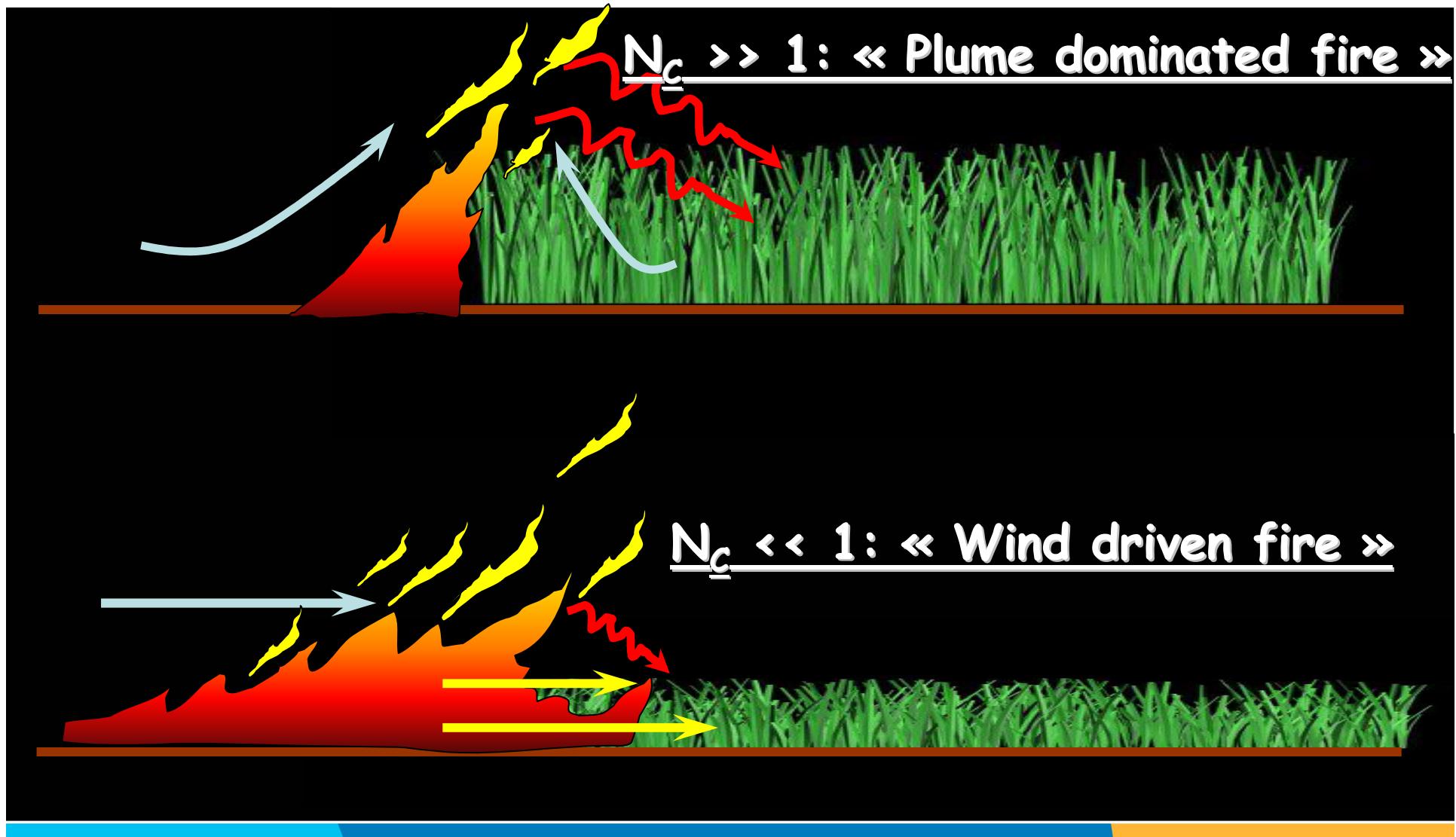
$$P_{Plume} = g \ I$$



$$P_{Wind} = \frac{1}{2} \rho C_P T_0 (U_W - ROS)^3$$

$$N_C = \frac{P_{Plume}}{P_{Wind}} = \frac{2 g I}{\rho C_P T_0 (U_W - ROS)^3}$$

# Wildfire propagation regimes



# Wildfires classification

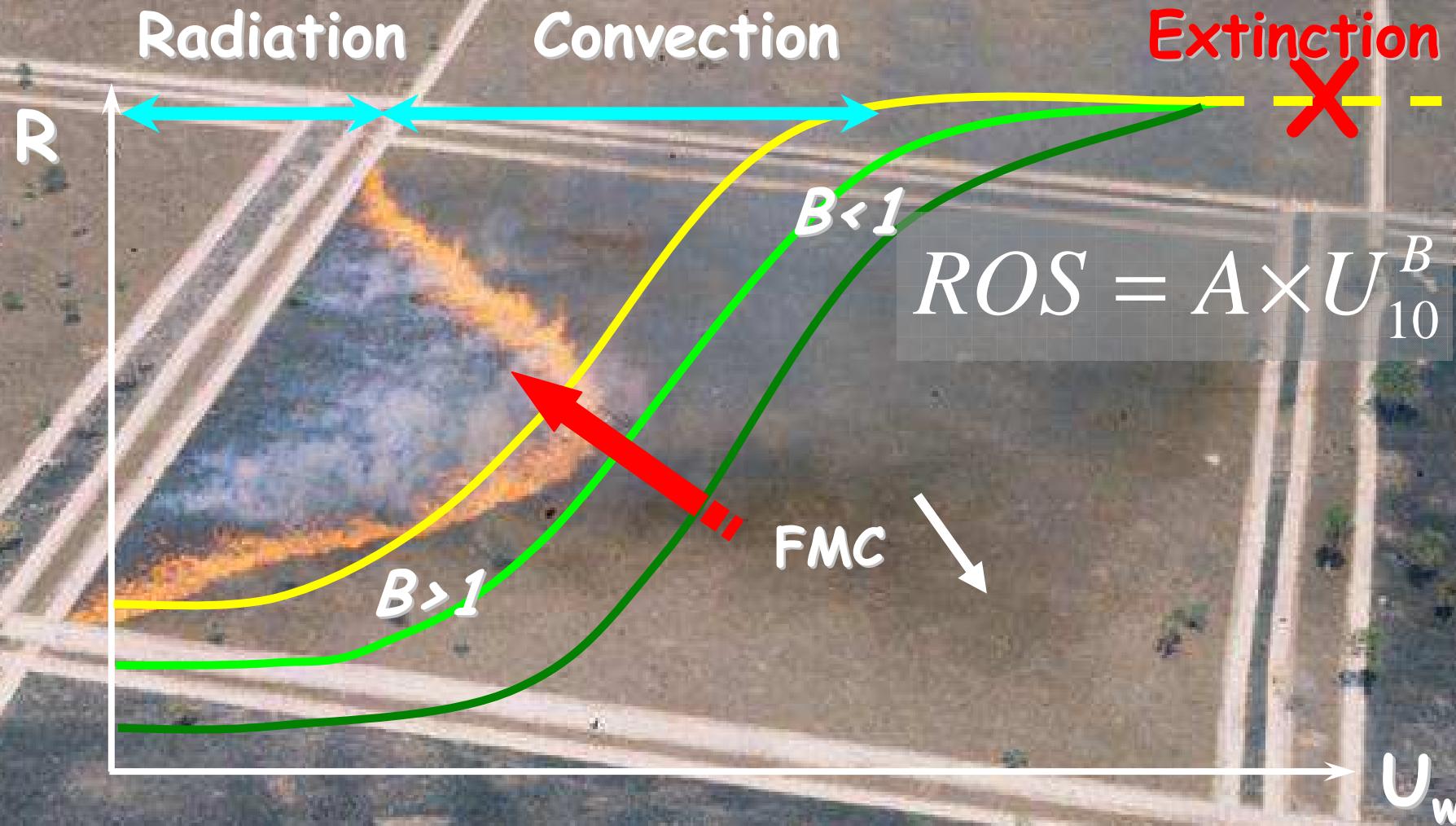
**Surface fires  
(wind driven)**



**Crown fires  
(plume dominated)**

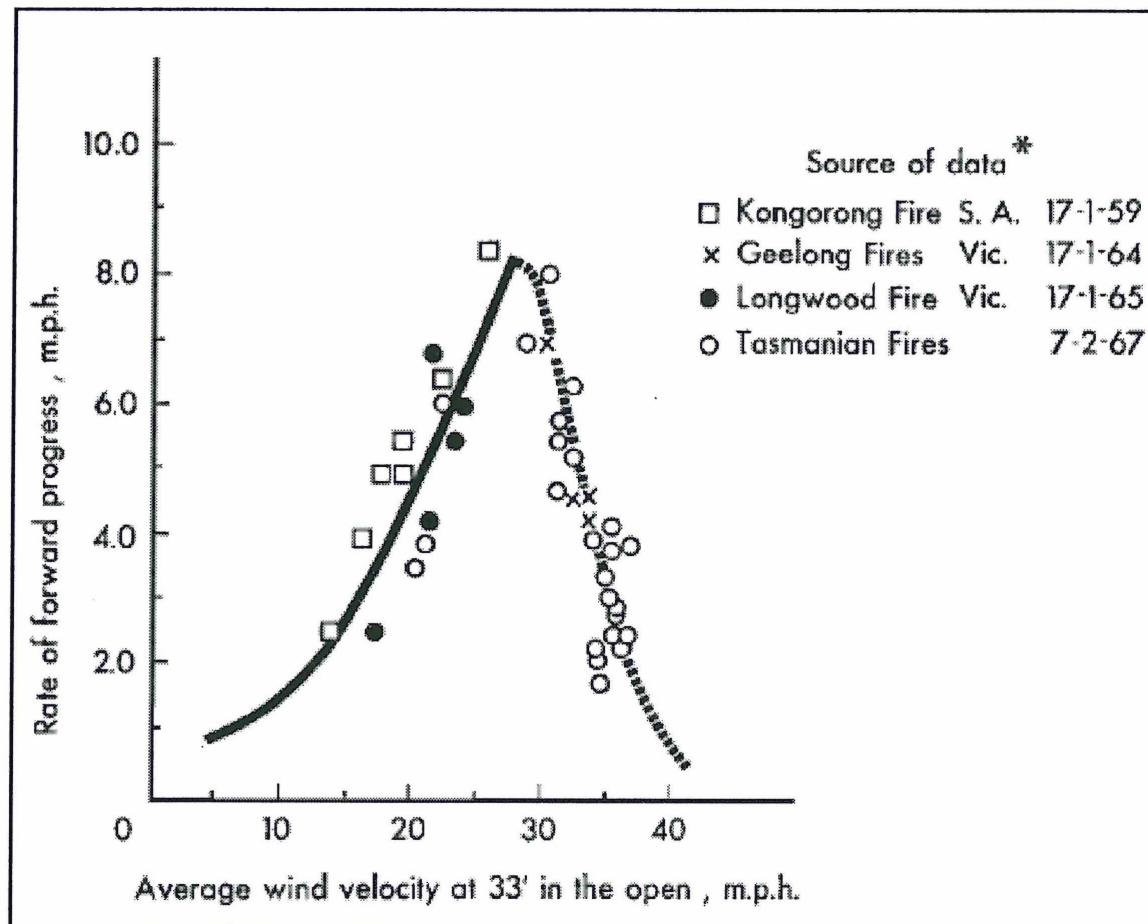


# Regimes of propagation of surface fires



Fogarty & al Fire Tech. Transfer Note 1999

# Rate of spread versus wind speed



McArthur 1969, Rothermel 1972

# Surface fires (grass)

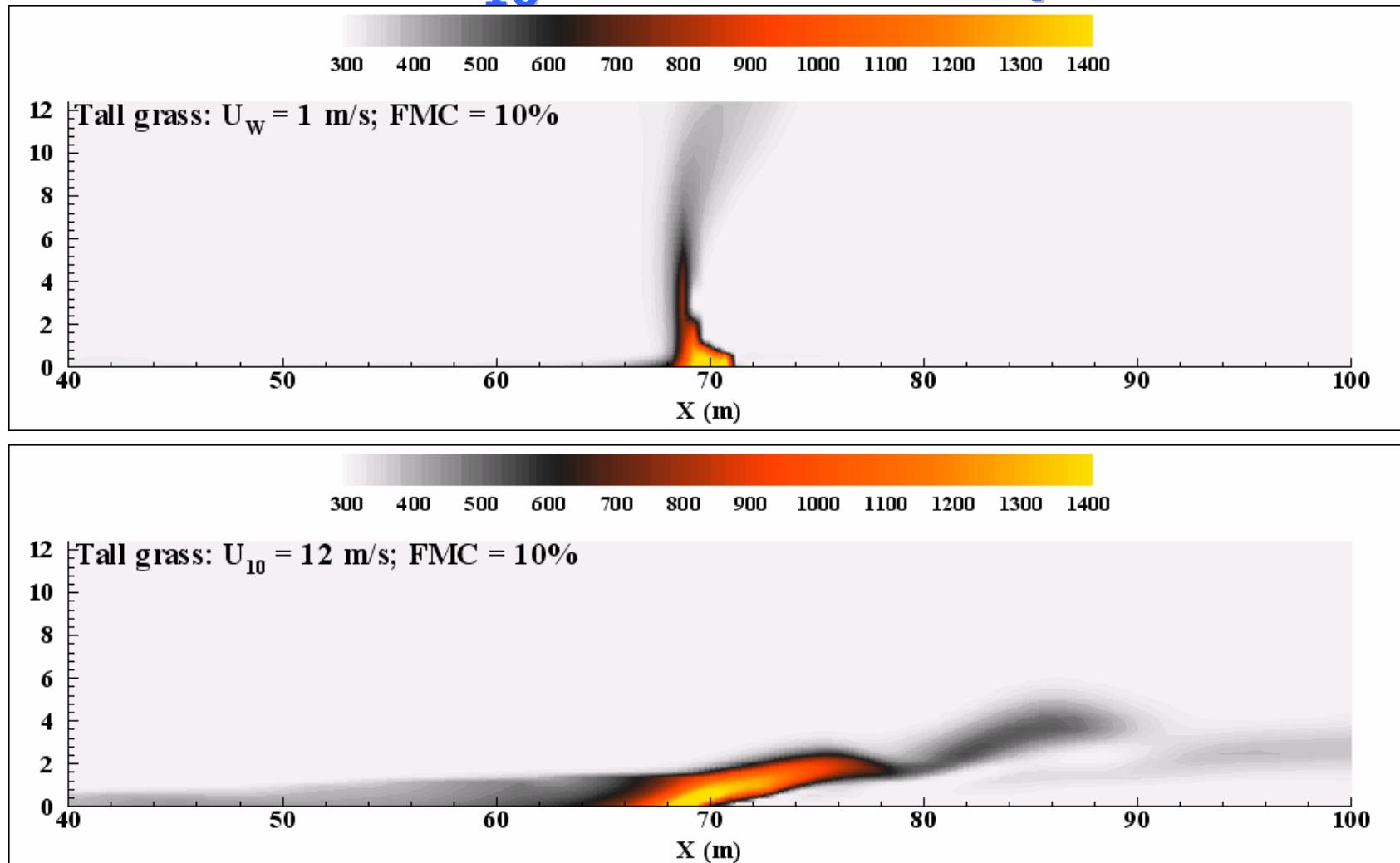
Fuel:

- Load = 7 t/ha
- Density = 500 kg/m<sup>3</sup>
- SAN = 4000 m<sup>-1</sup>
- FMC = 10 %
- Wind speed = 1 – 25 m/s
- Nc ~ 0.1 - 1400

Cheney & al IJWF 1993

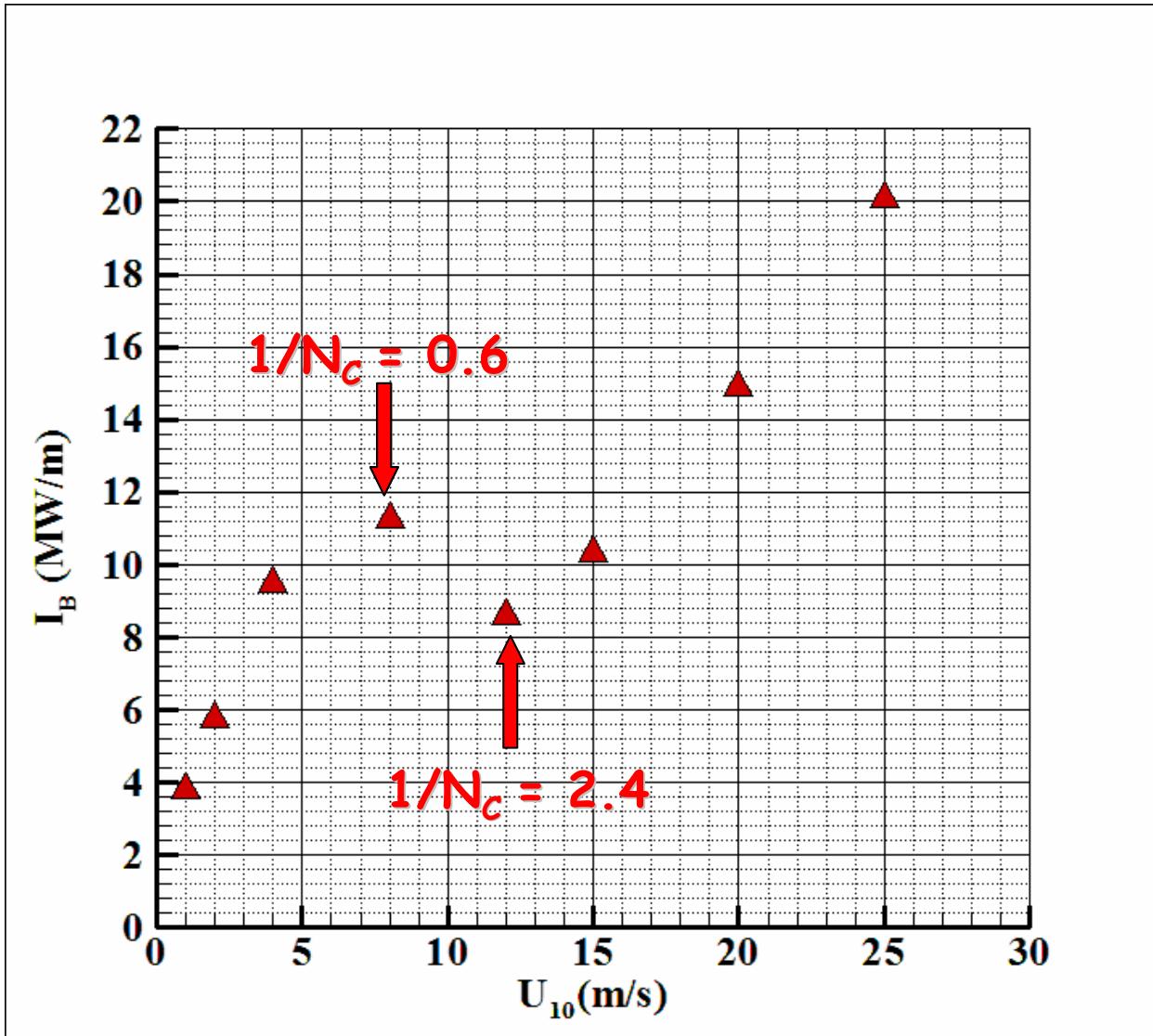


# Grassland fires: temperature field for $U_{10} = 1$ and 12 m/s

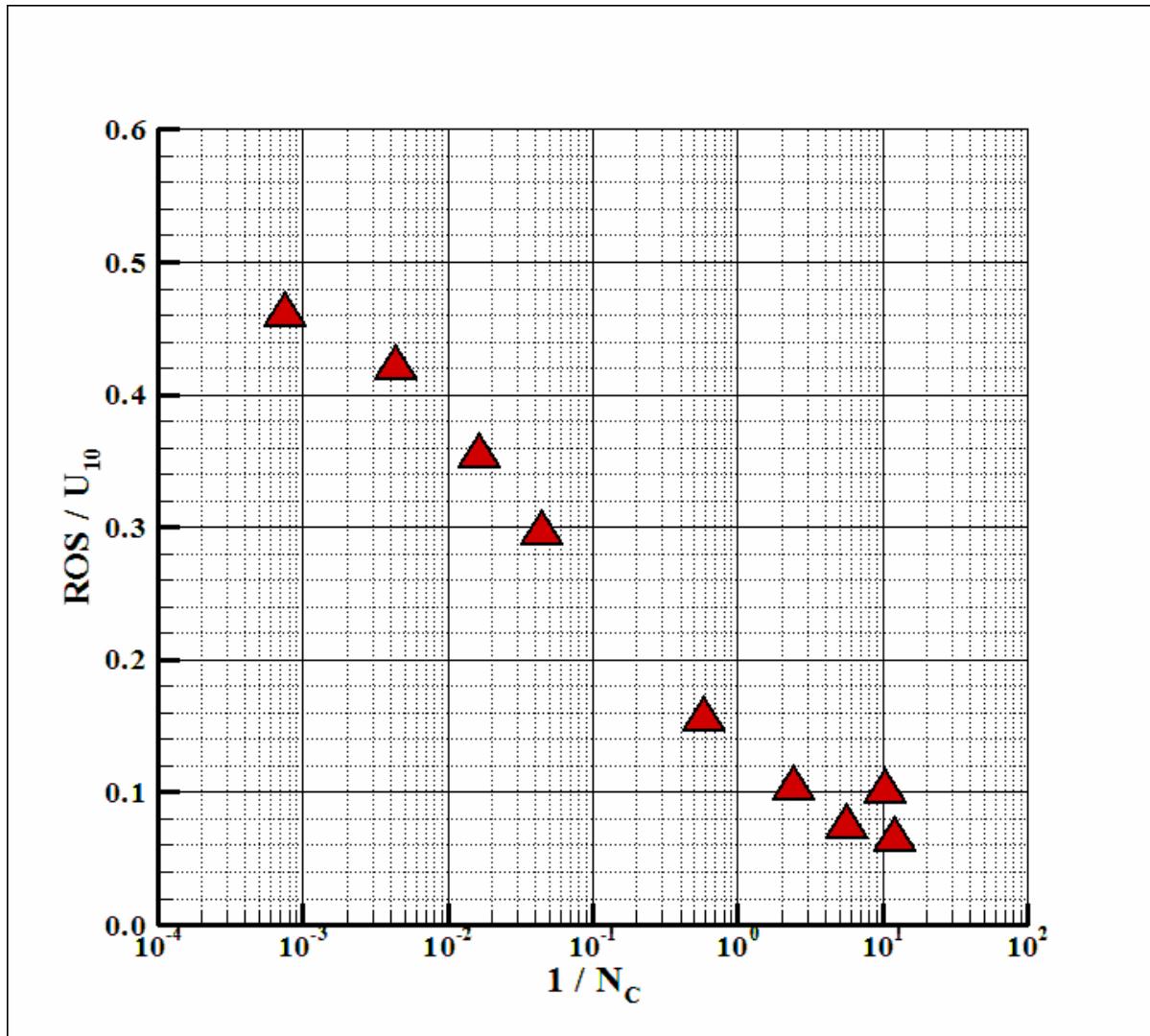


(Morvan & al FSJ 2009, Morvan CST 2014)

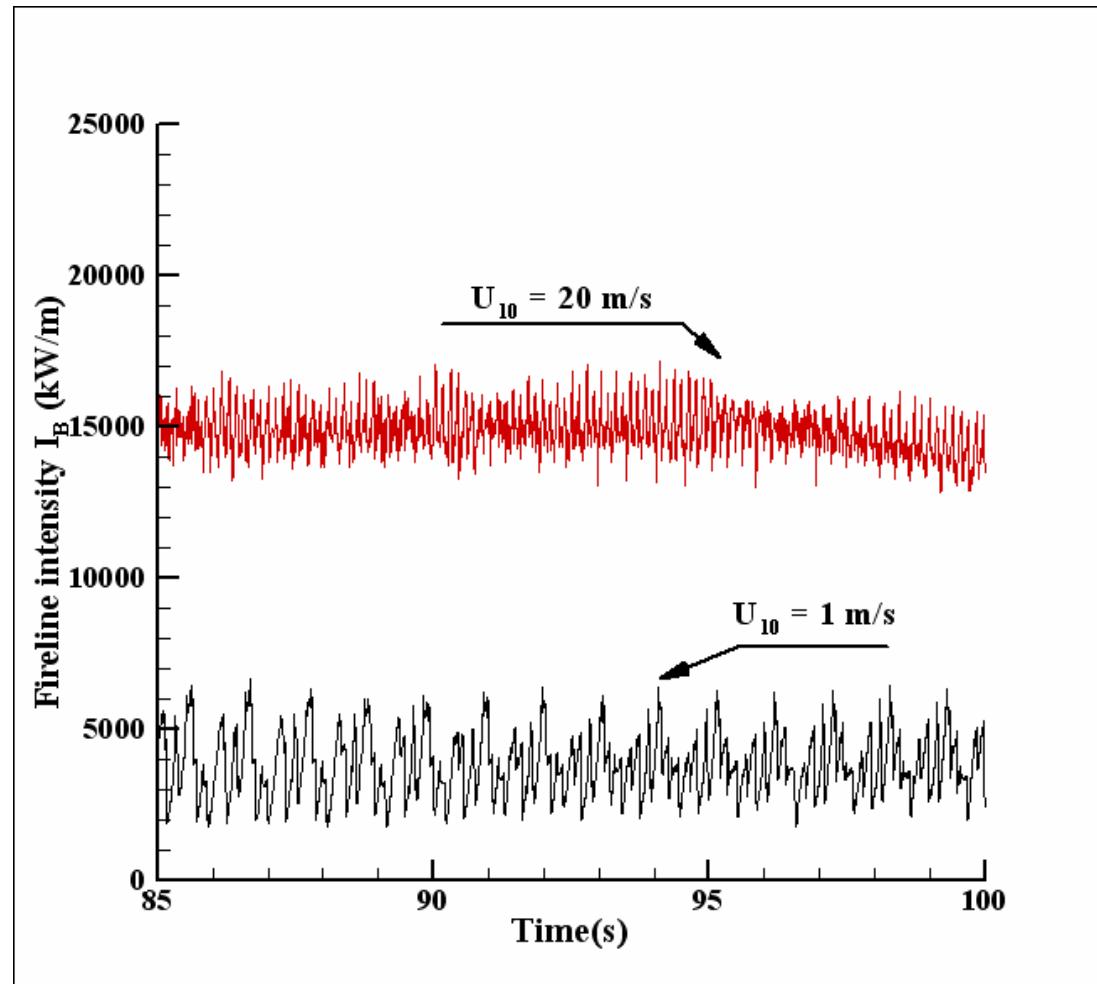
# Fireline intensity versus wind speed



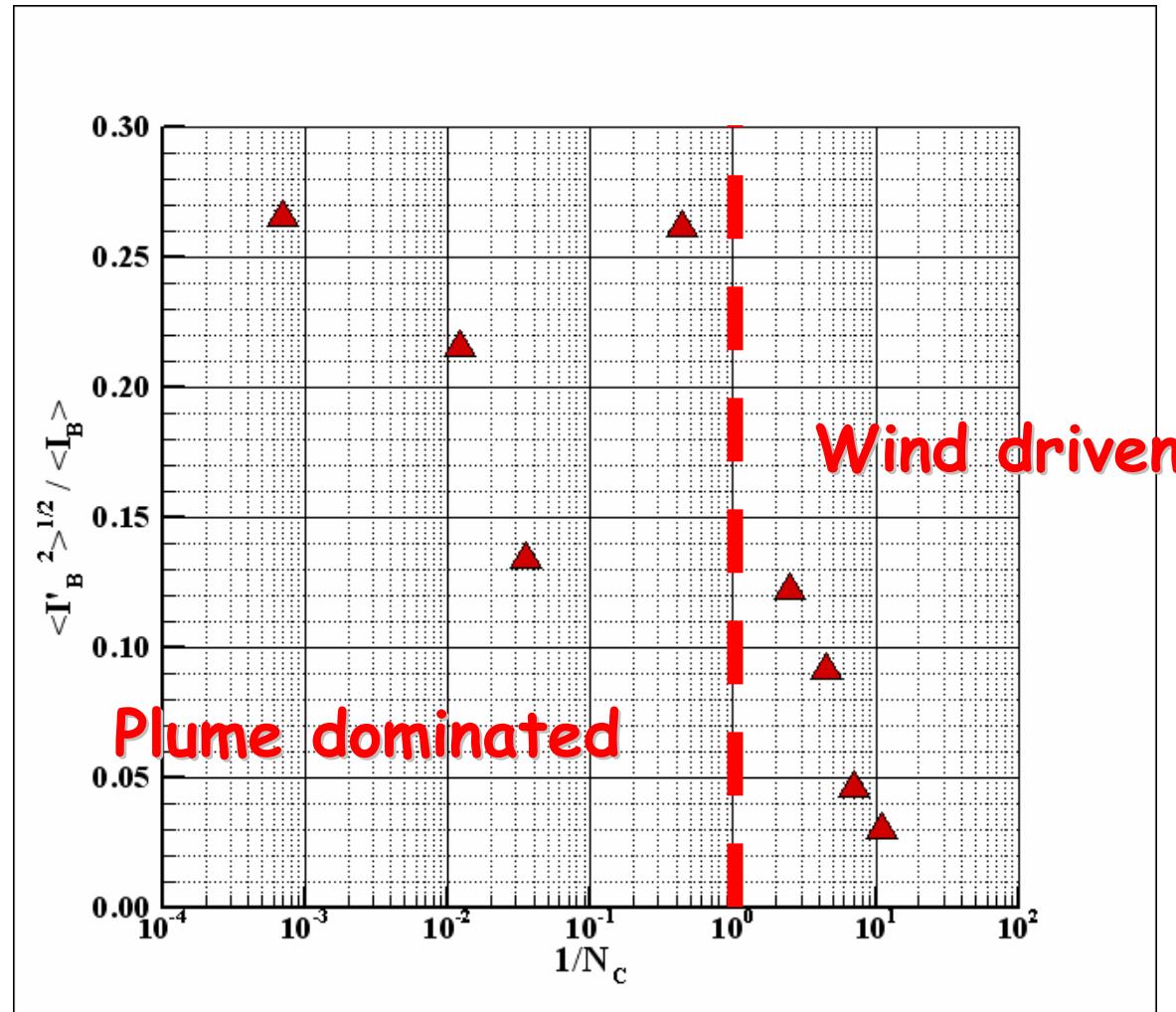
## Convective Byram number / regime of propagation



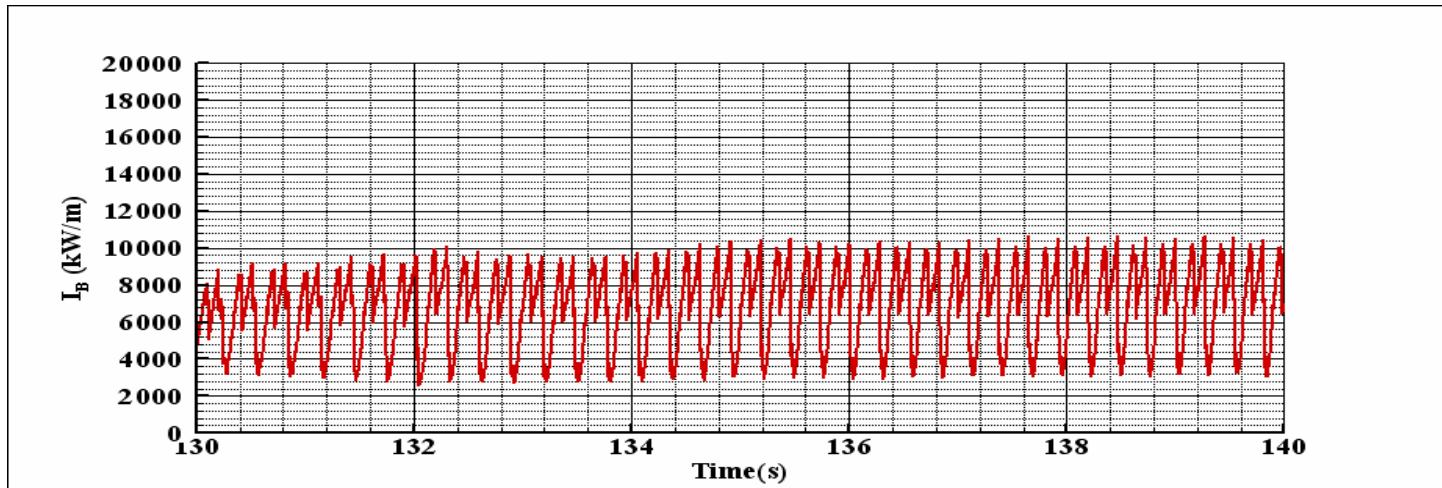
## Time evolution of the fireline intensity



# Fireline intensity (normalized standard deviation) versus convective Byram number

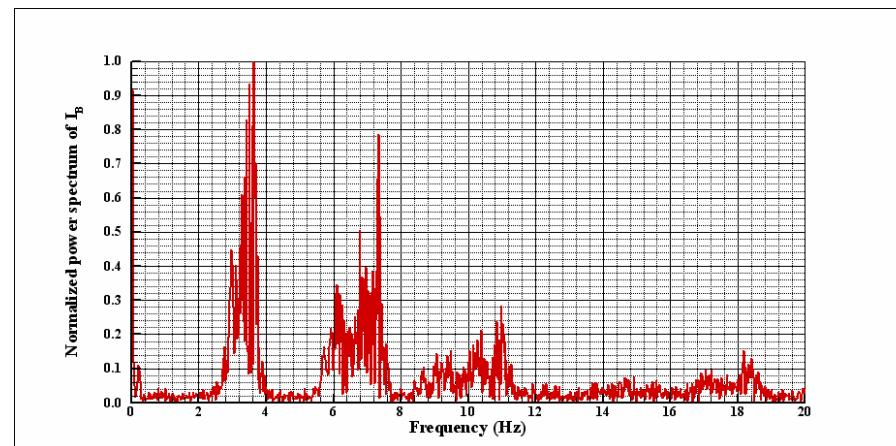


## Time history and spectrum of fire intensity ( $U_{10} = 2 \text{ m/s}$ )



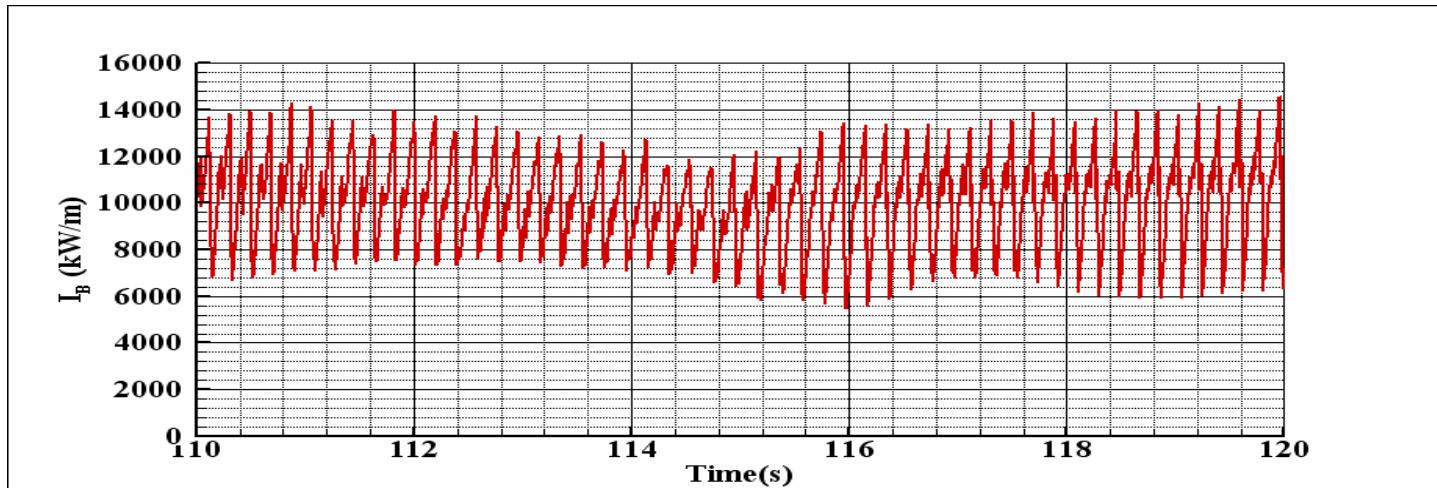
$$f_{KH} = 0.26 \text{ Hz}$$

$$f_B = 1.3 \text{ Hz}$$

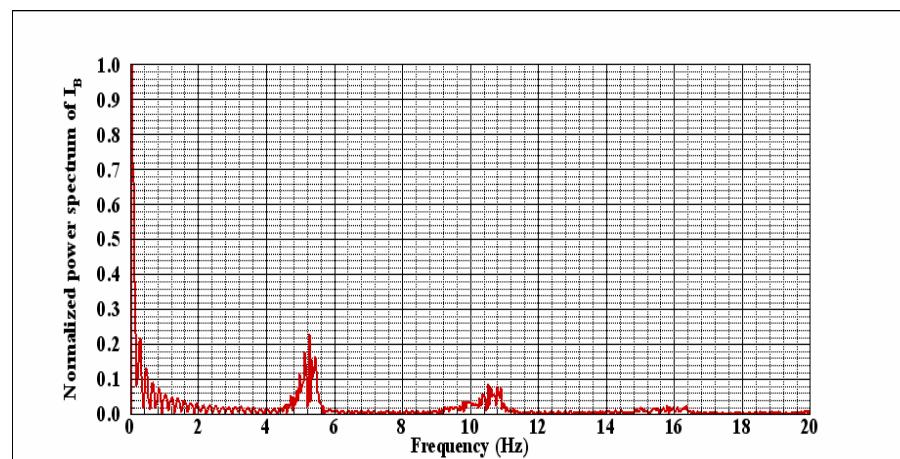


$$f = 3.4 \text{ Hz}$$

## Time history and spectrum of fire intensity ( $U_{10} = 20 \text{ m/s}$ )



$$f_{KH} = 2.66 \text{ Hz}$$
$$f_B = 0.66 \text{ Hz}$$

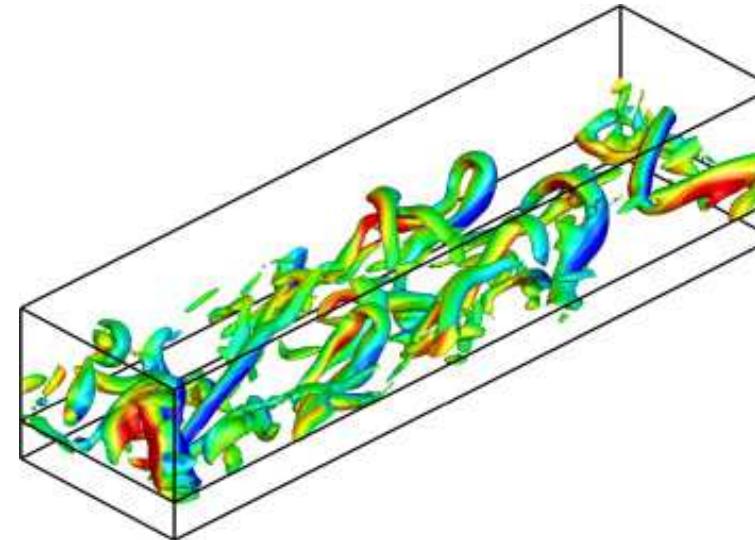


$$f = 5.5 \text{ Hz}$$

# Instabilities associated with a fire front

$$S_t^{KH} = \frac{f_{KH} H_{Fuel}}{U_{10}} = 0.093$$

$f_{KH} \nearrow$  if  $U_H \nearrow$

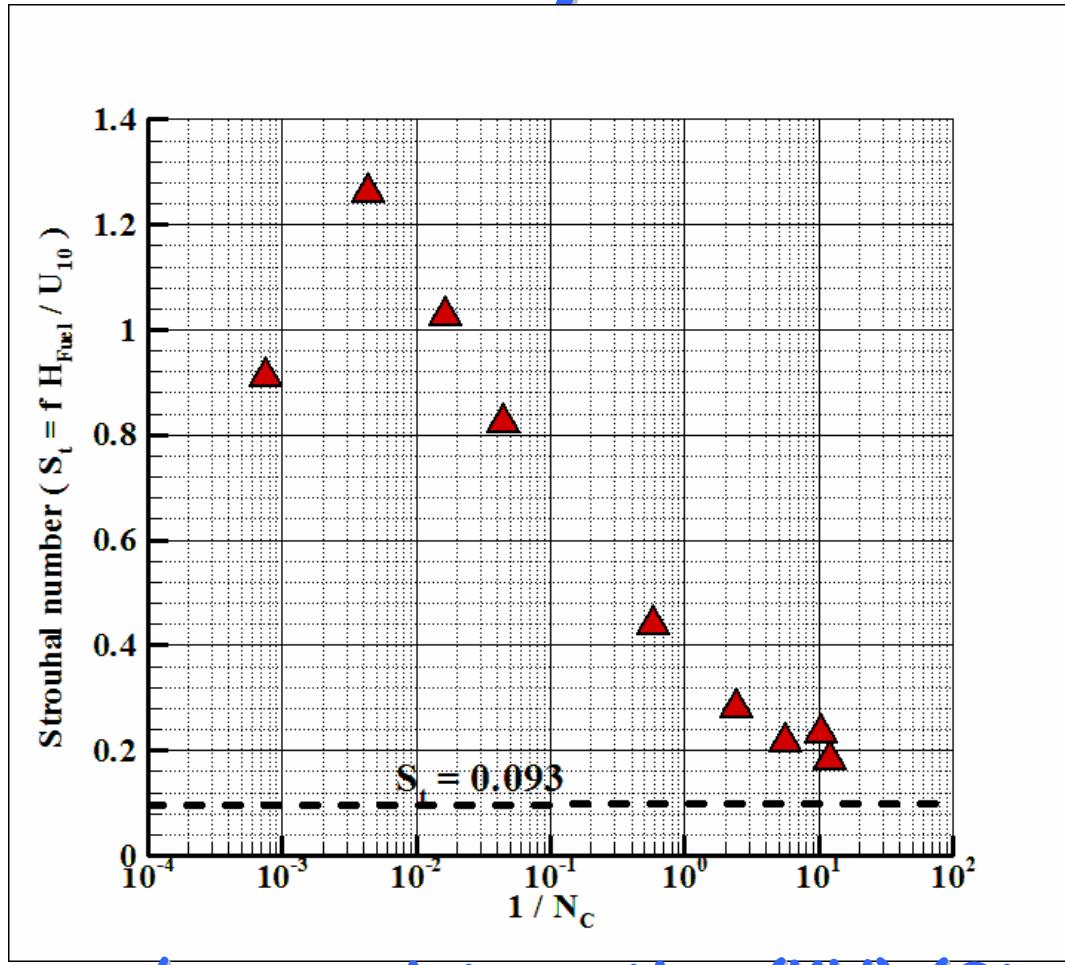


$$S_t^B = f_B \sqrt{\frac{D_{Fire}}{g}} \approx 0.5$$

$D_{Fire} \nearrow$  if  $U_H \nearrow \Rightarrow f_B \searrow$

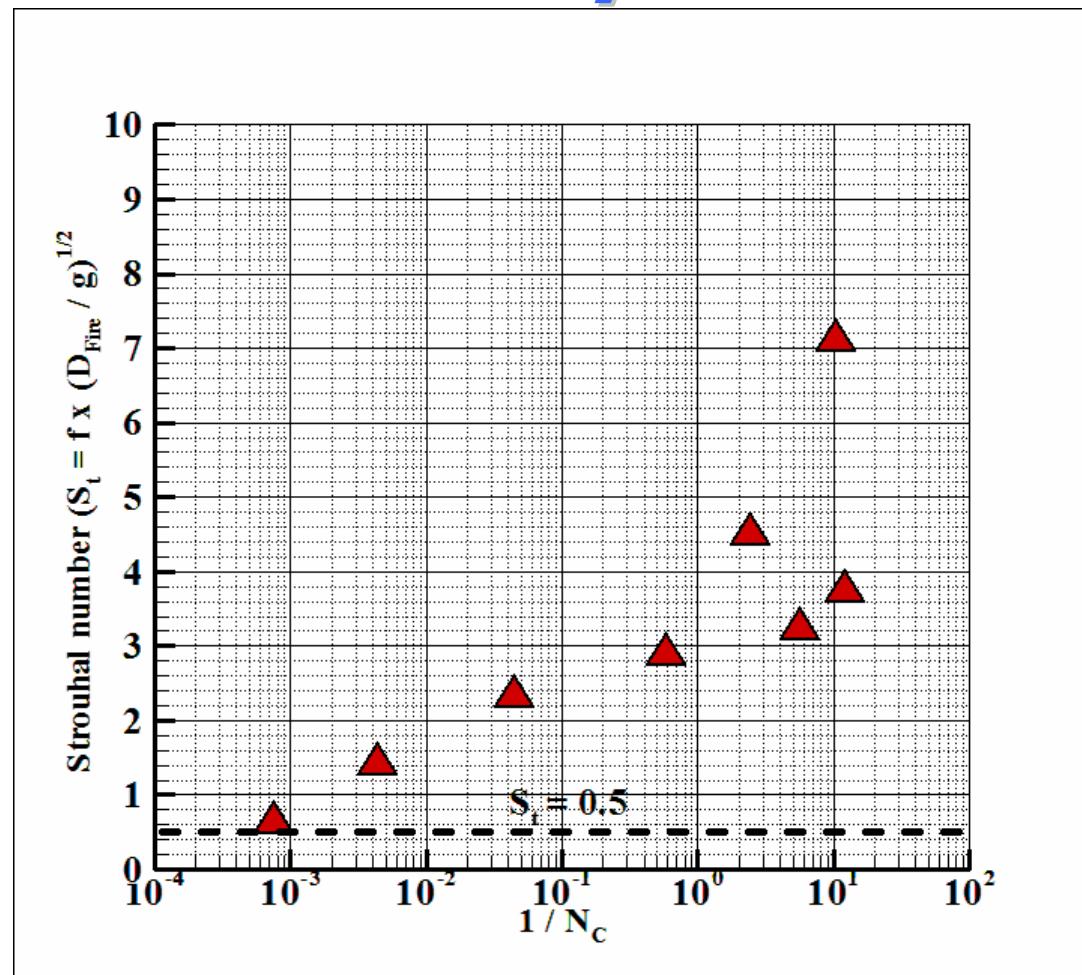


# Strouhal number (wind scale) versus convective Byram number



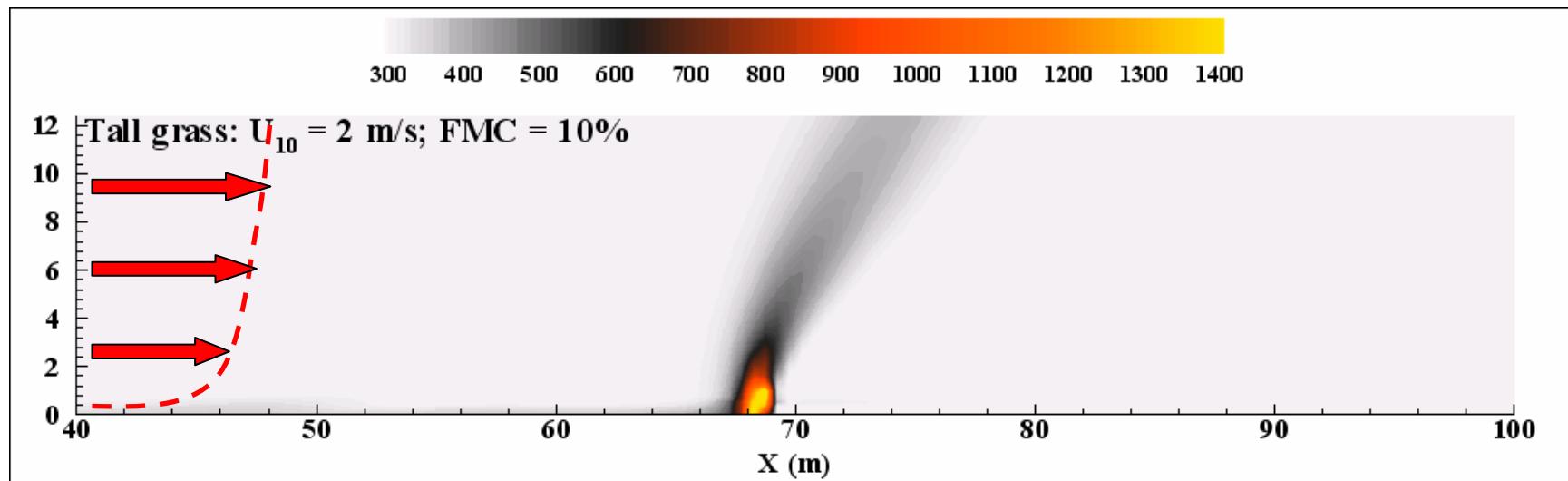
Turbulence/canopy interaction (KH) ( $St = 0.093$ )  
(streamwise direction)

# Strouhal number (plume scale) versus convective Byram number



Pool fire ( $St = 0.5$ )

## Effect of an unsteady (sinusoidal) inlet wind flow

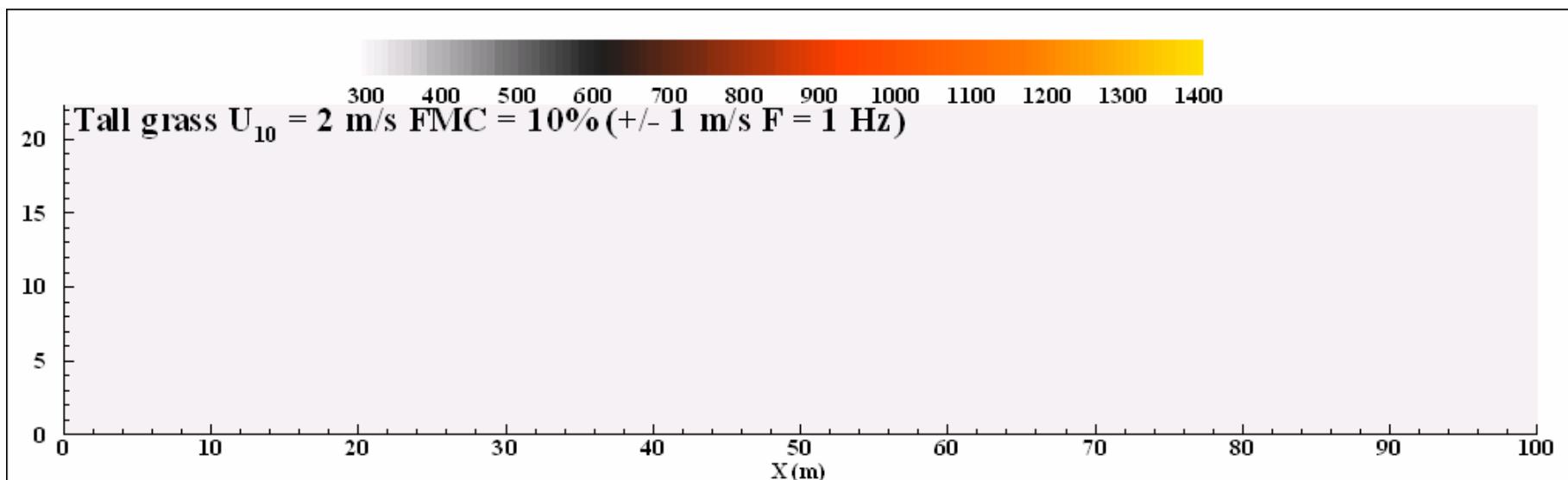
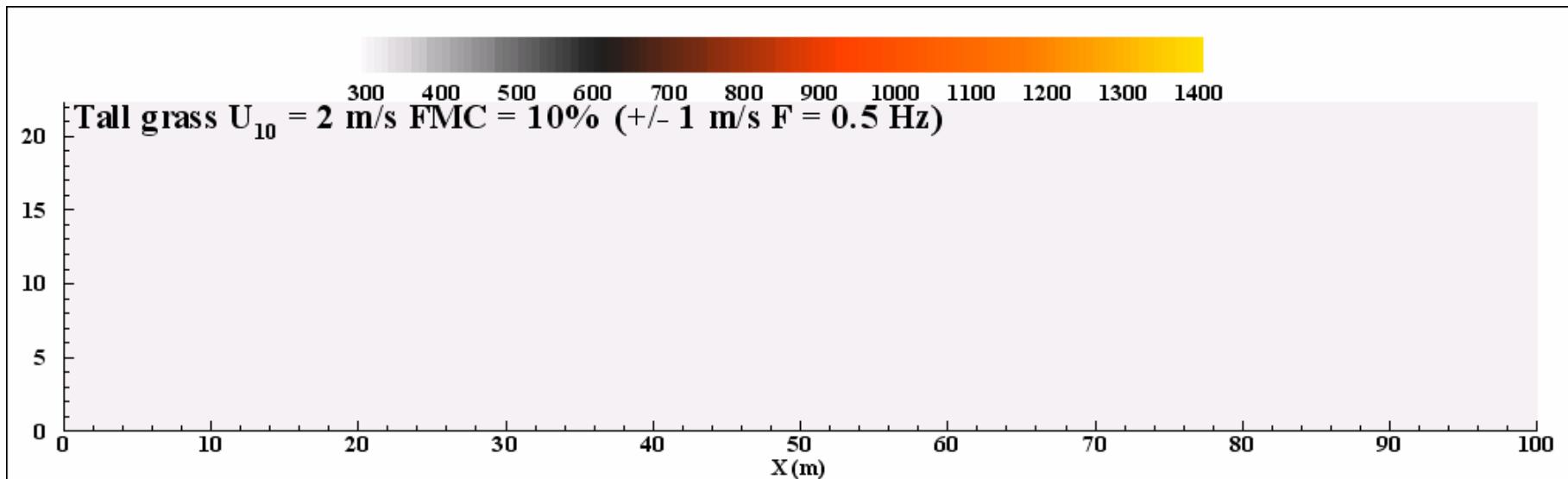


$$U_x = A \times U_{10} \times \ln\left(\frac{z + z_0}{z_0}\right) + \Delta U_{10} \sin(2\pi \times f \times t)$$

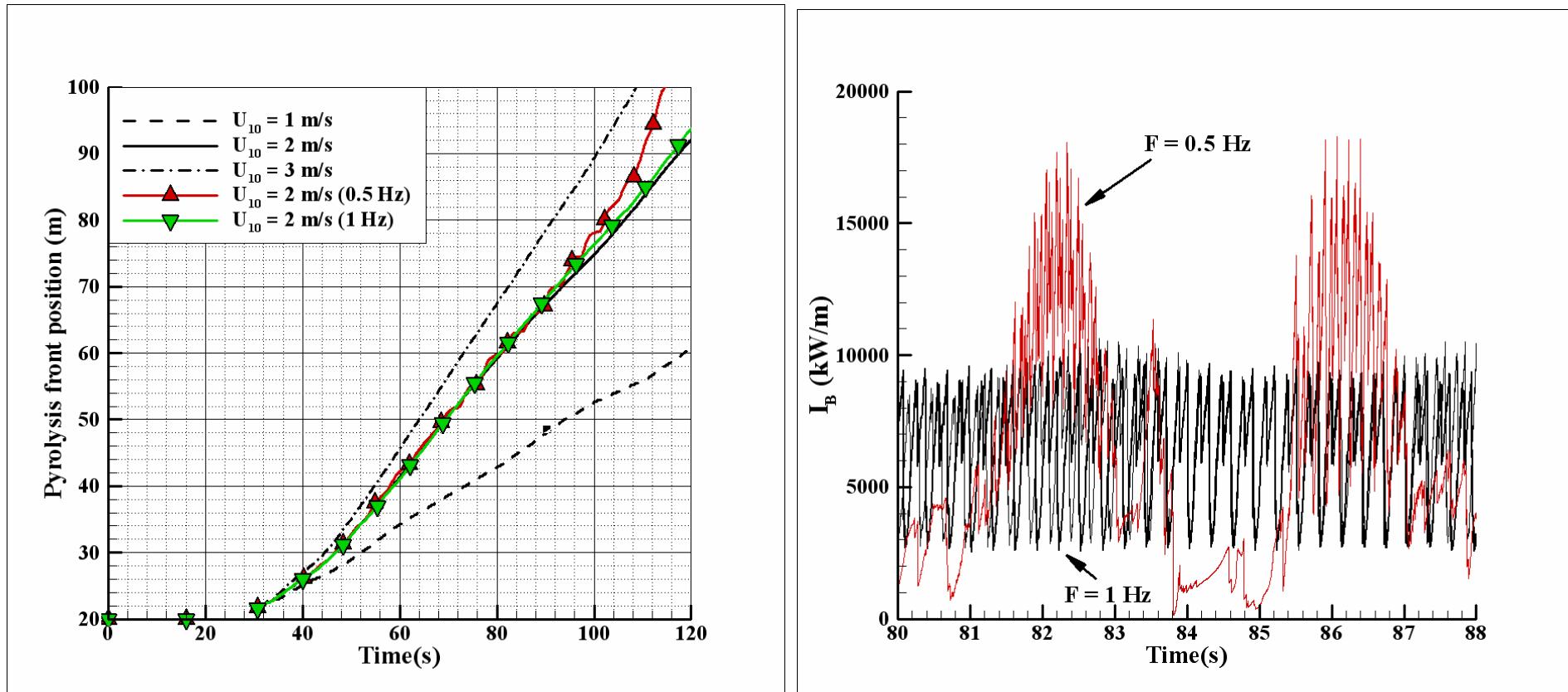
$$U_{10} = 2 \text{ m/s} \quad \Delta U_{10} = 1 \text{ m/s} \quad f = 0 - 3 \text{ Hz} \quad z_0 = 0.01 \text{ m}$$

Pitts Prog. Energy Combust. Sci. 1991  
Morvan Combust. Sc. Tech. 2014

## Grassland fires: temperature field $U_{10} = 2 \text{ m/s} +/- 1 \text{ m/s}$



# Trajectory of the pyrolysis front ( $T_s = 500$ K) Fireline intensity

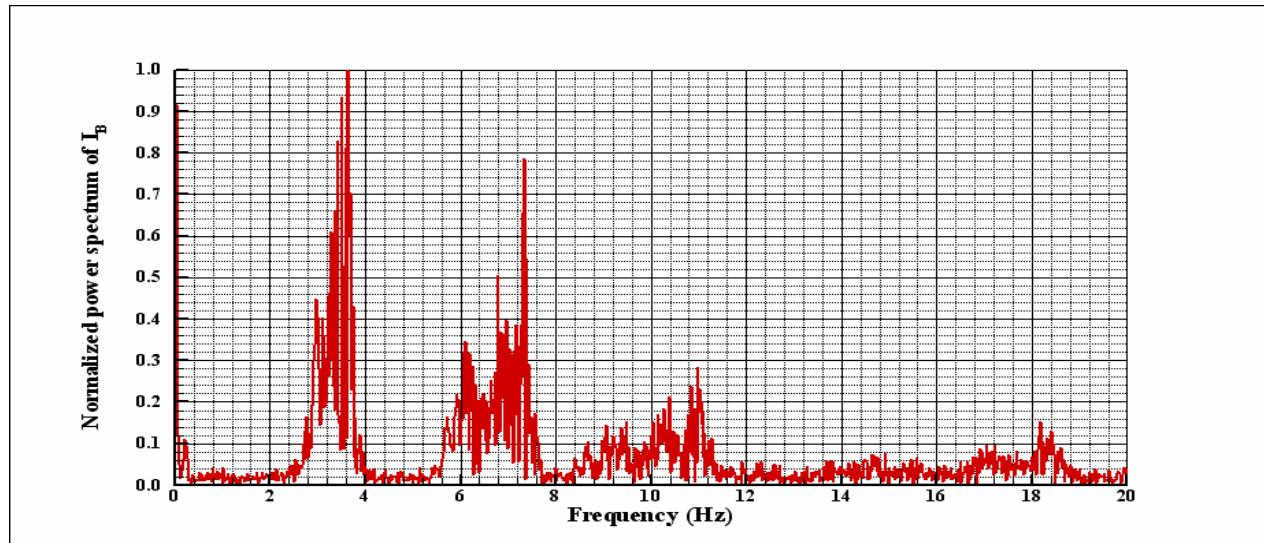


**Steady  $\rightarrow \langle I_B \rangle = 6505$  kW/m**

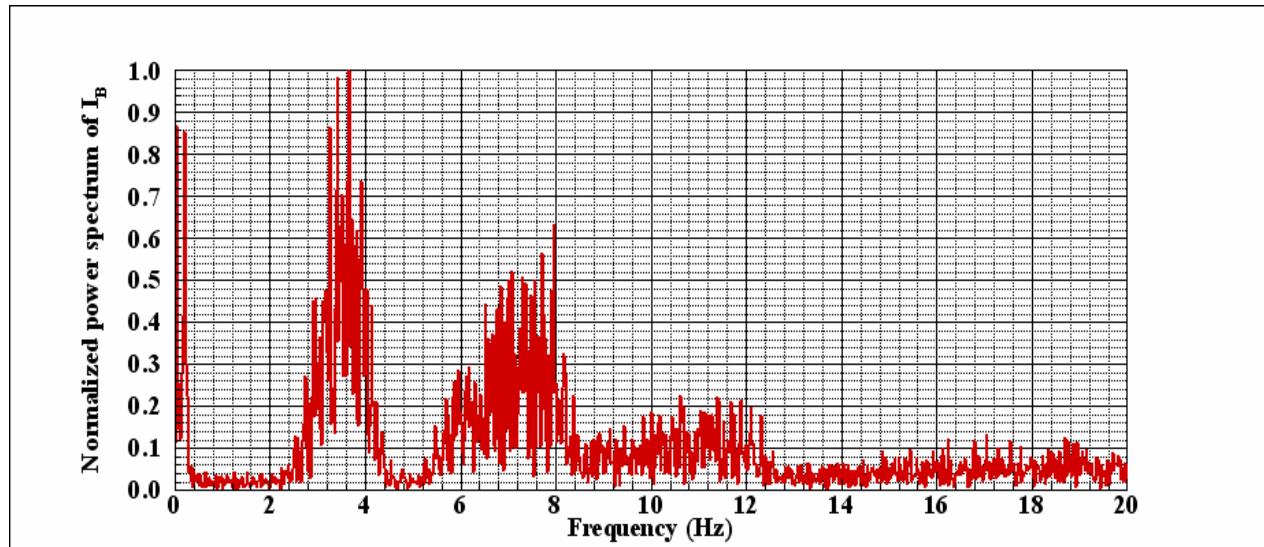
$F = 0.5$  Hz  $\rightarrow \langle I_B \rangle = 6972$  kW/m (+7%)

$F = 1$  Hz  $\rightarrow \langle I_B \rangle = 6742$  kW/m (+3.6%)

# Fireline intensity normalized spectrum

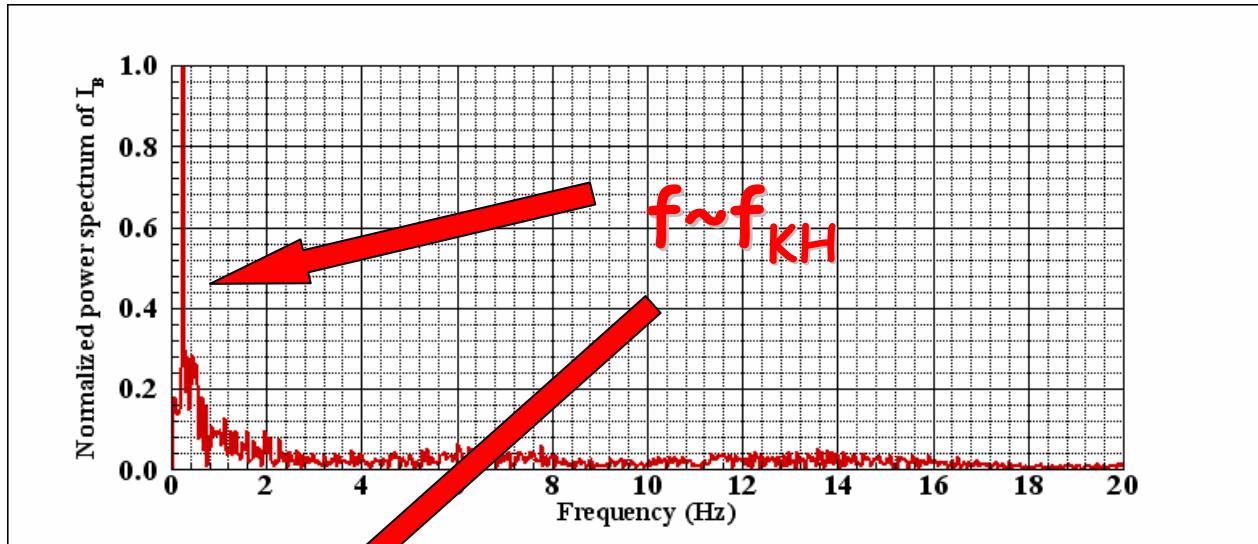


steady

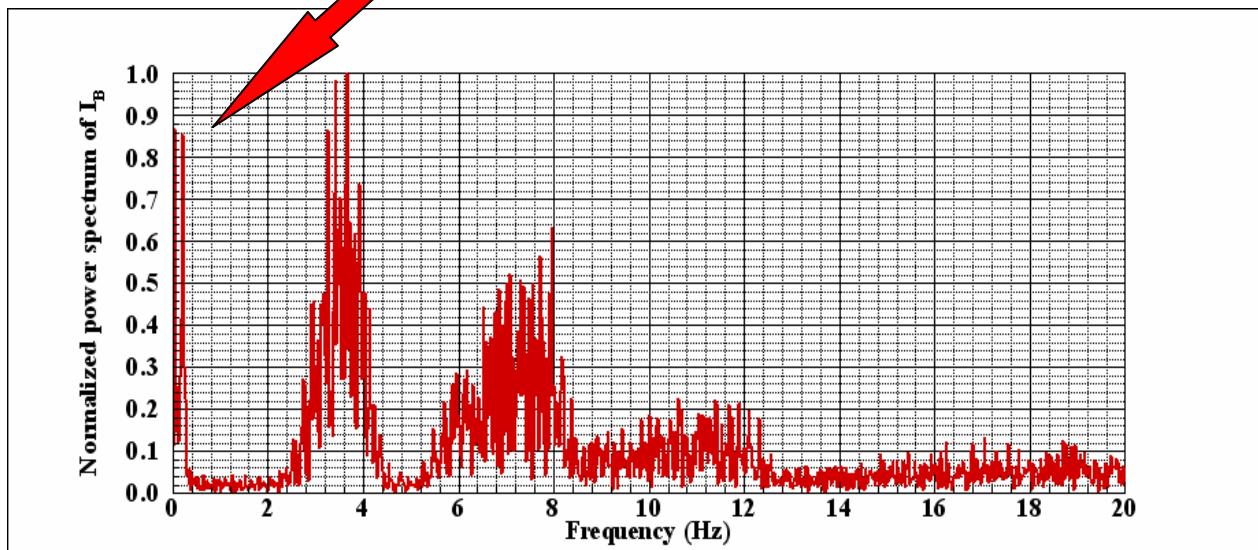


F=1 Hz

# Fireline intensity normalized spectrum



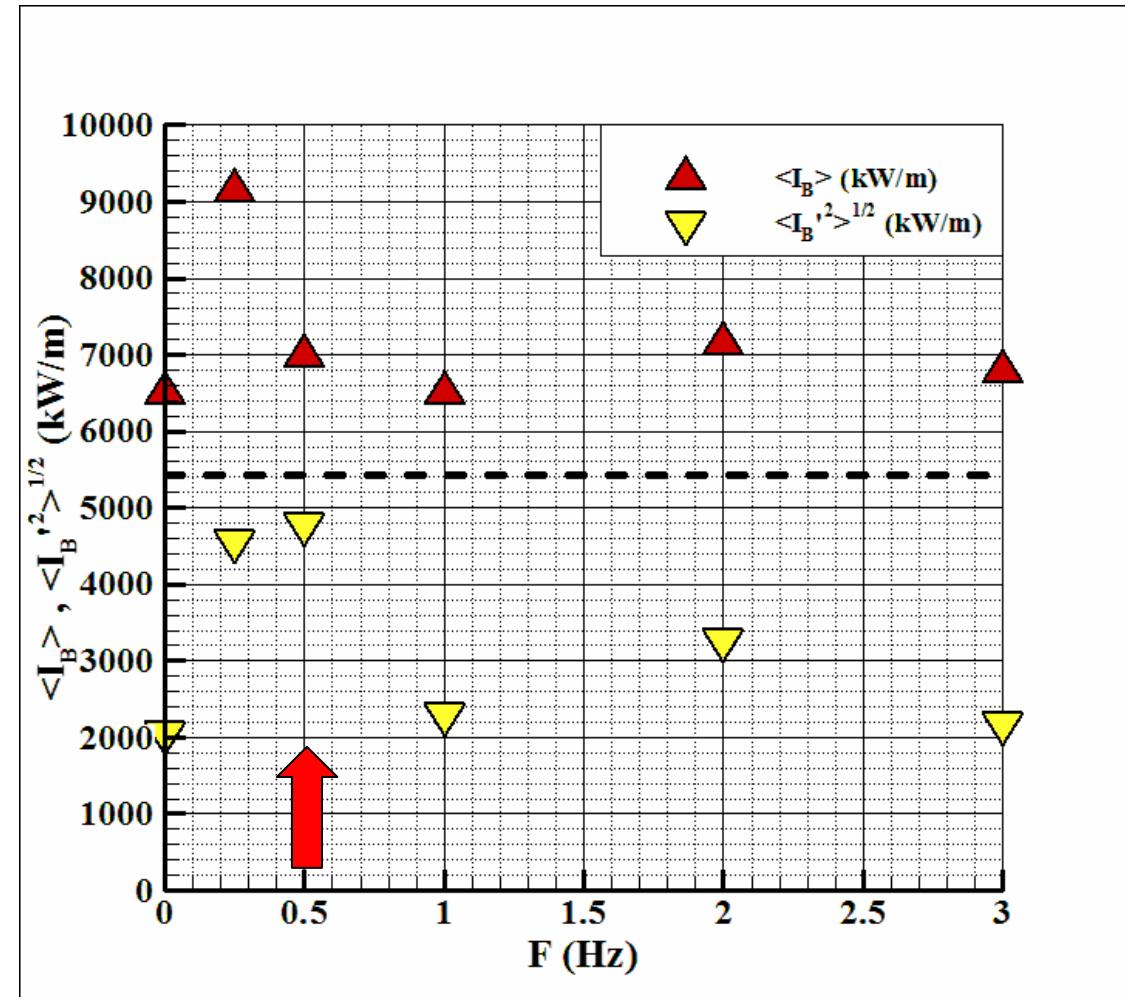
$F = 0.5 \text{ Hz}$



$F = 1 \text{ Hz}$

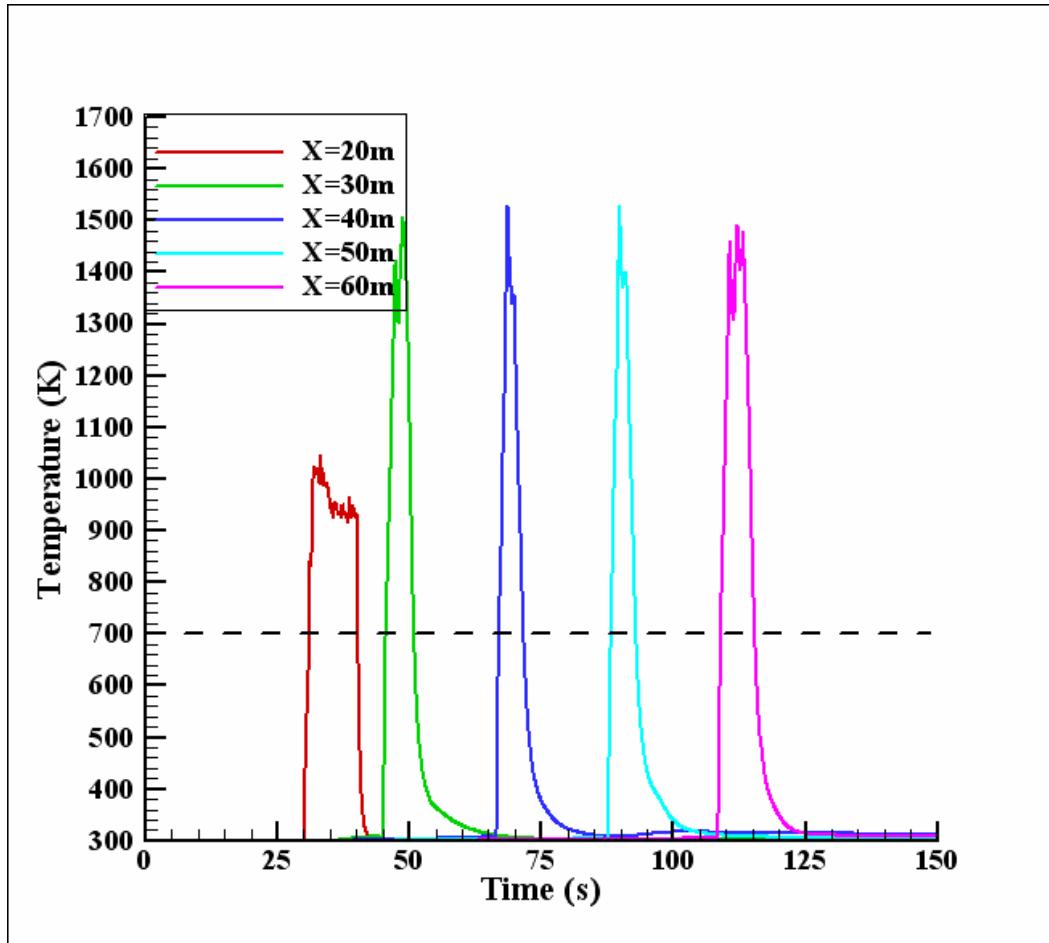
## Fire intensity and standard obtained for $U_{10} = 2 \text{ m/s}$ and a sinusoidal time variation ( $f = 0.26 - 3 \text{ Hz}$ )

$(f_{KH} = 0.26 \text{ Hz})$



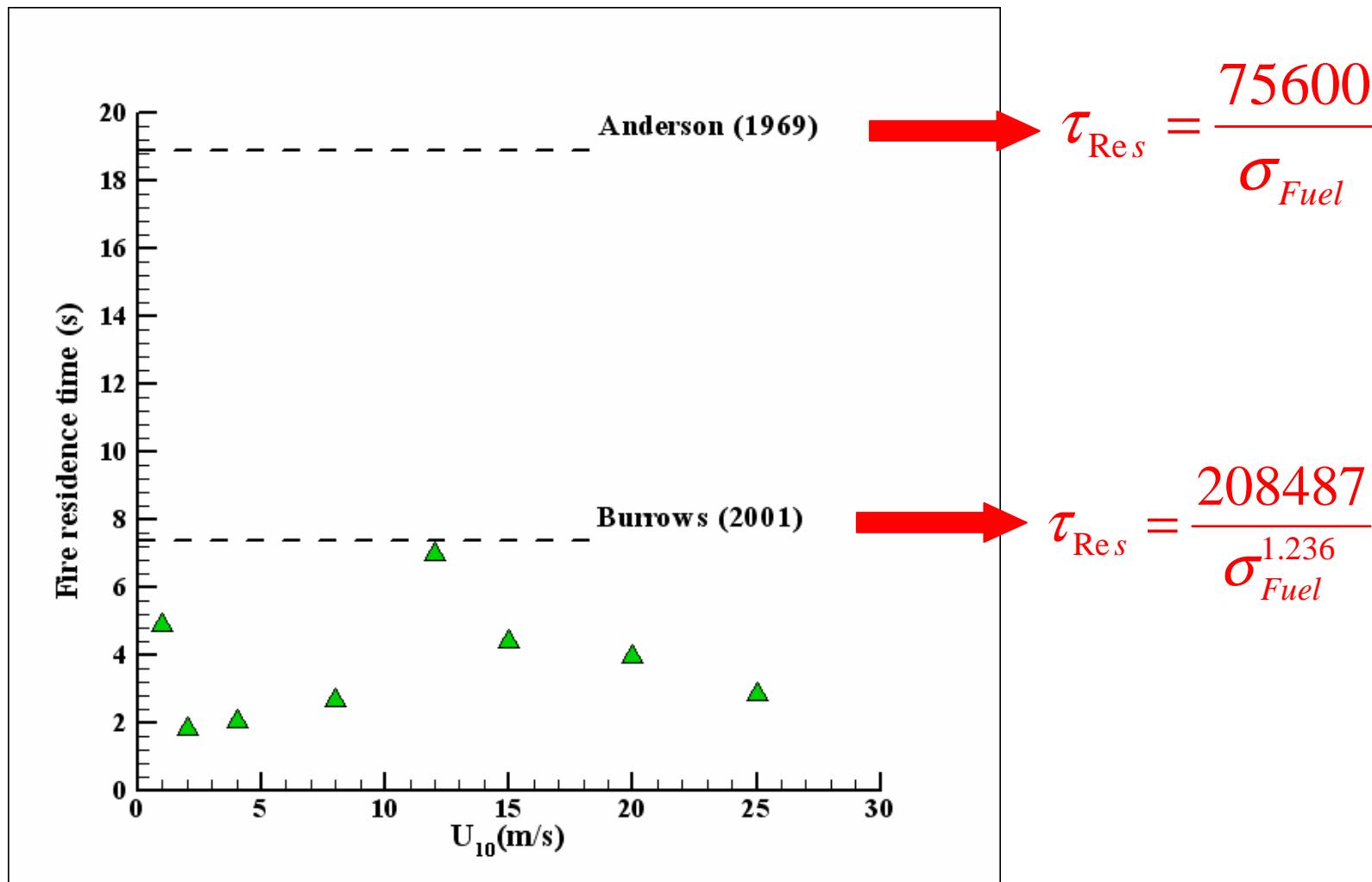
$2 \times f_{KH}$

# Evaluation of the fire residence time from the time history of the temperature in the fuel layer

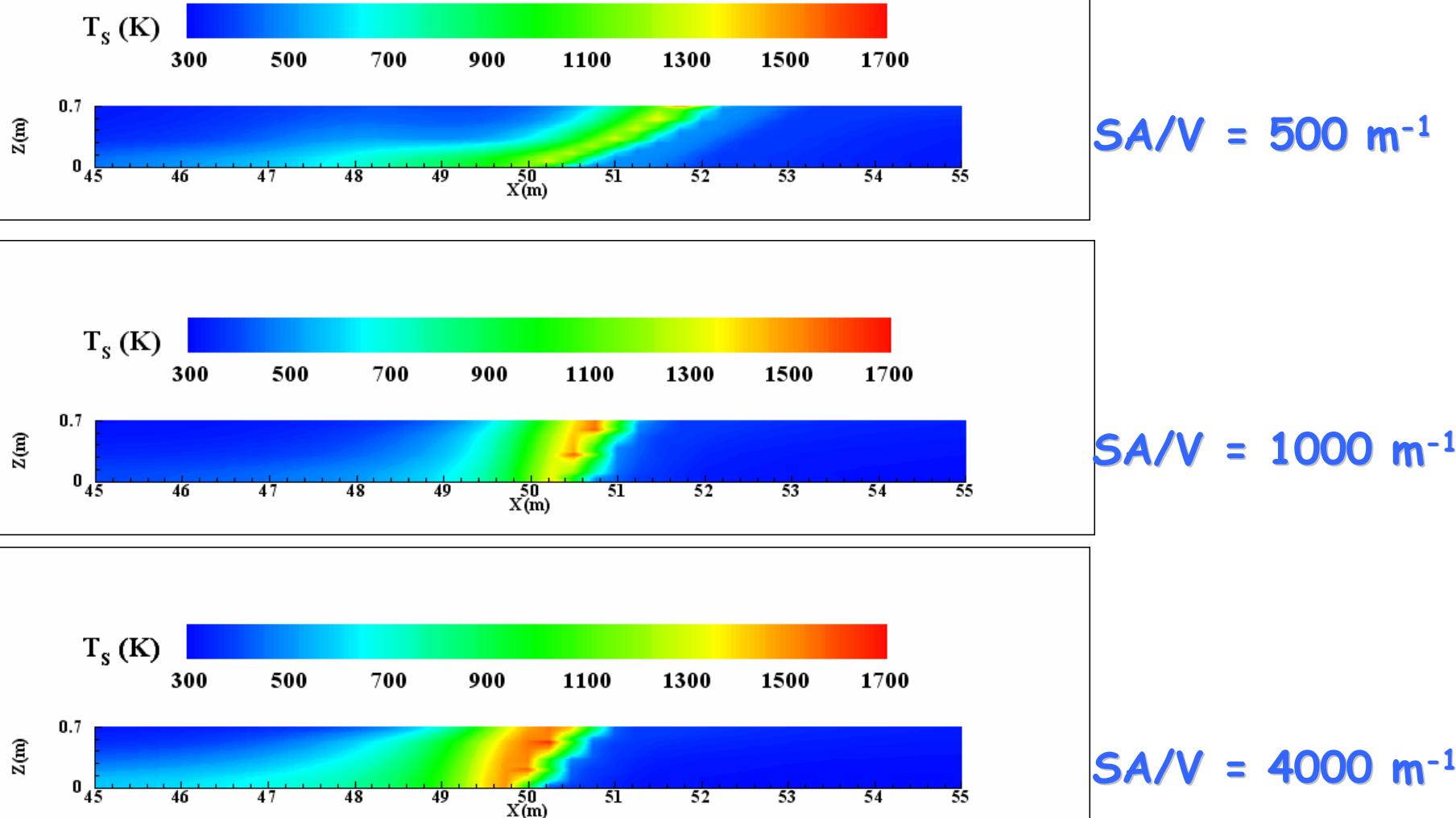


( $z = 0.25 \text{ m}$ , fuel depth = 0.7 m)

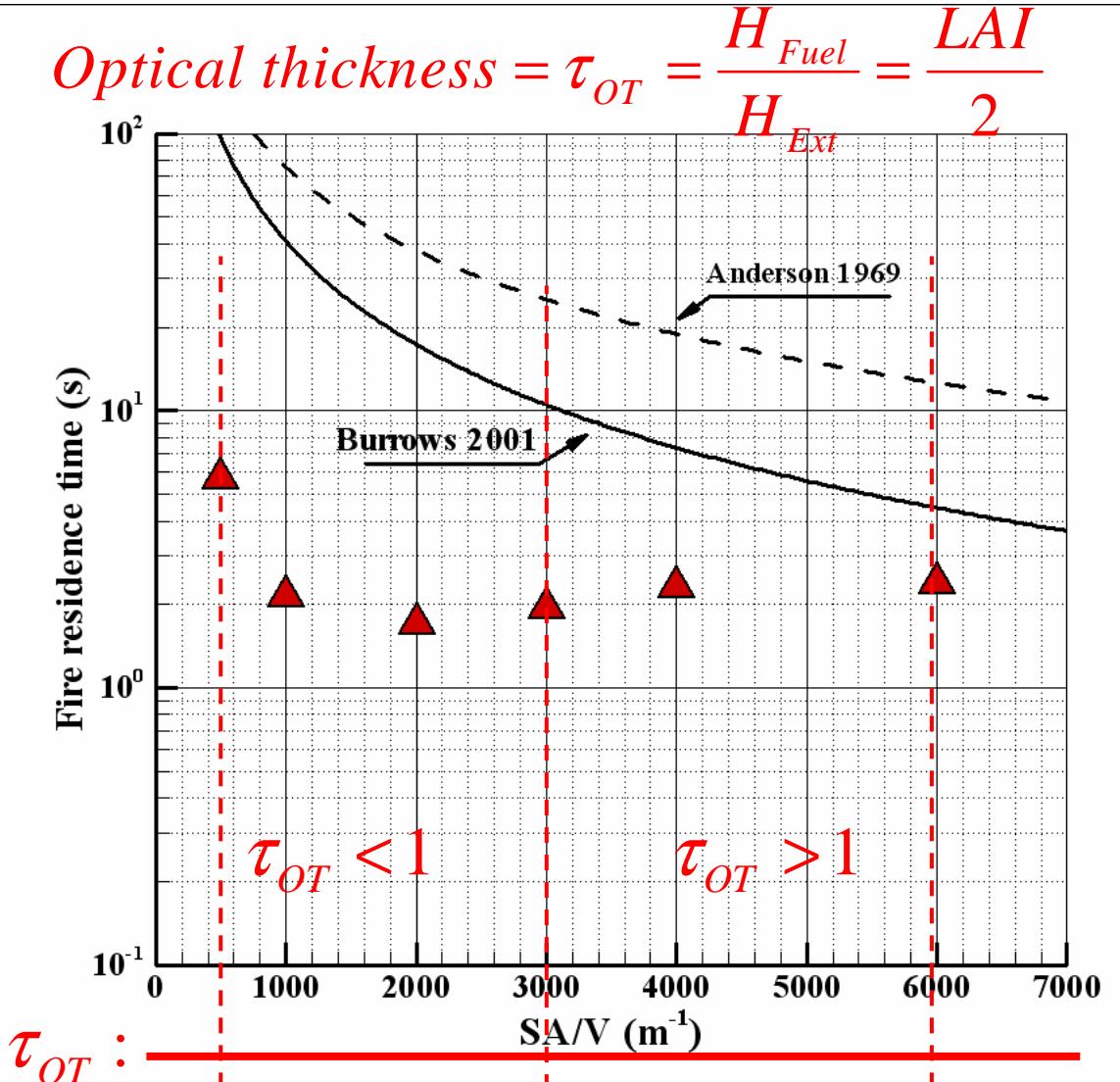
## Fire residence time versus 10m open wind velocity



## Solid fuel temperature field for various SA/V ( $U_{10} = 4 \text{ m/s}$ , FMC = 10%)



# Fire residence time versus SA/V



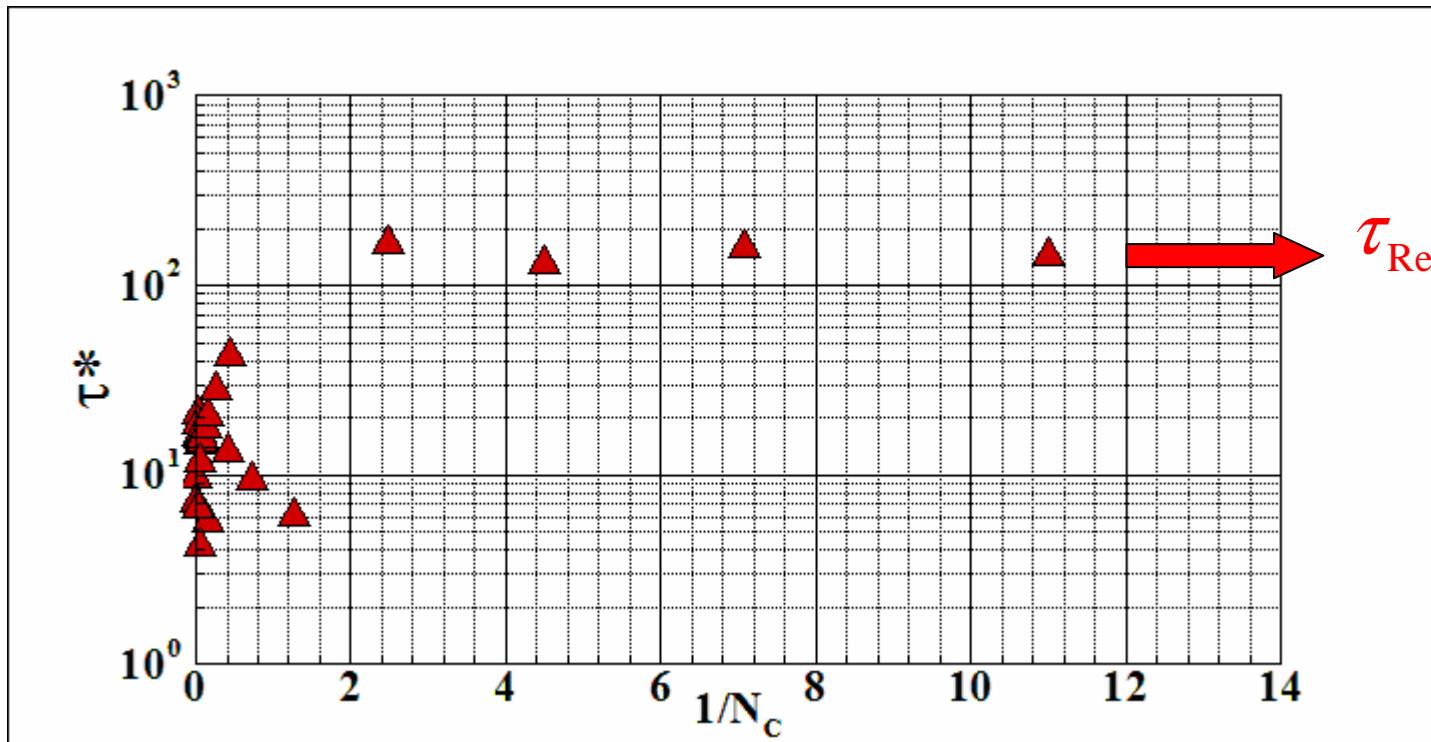
0.175

1.05

2.1

# Reduced fire residence time versus Byram convective number

$$\tau_{\text{Res}}^* = \frac{\tau_{\text{Res}} \times U_{10} \times LAD}{2} = \frac{\tau_{\text{Res}} \times U_{10} \times \alpha_{\text{Fuel}} \times \sigma_{\text{Fuel}}}{4}$$



$$= \frac{75000}{\sigma_{\text{Fuel}}}$$

# Thank you for attention Questions ?

