### **Climate Change Risk Assessment**

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Workshop on Uncertainty Quantification in Climate Modeling and Projection 2015 July 13-17







an NSF-sponsored research network for **Sustainable Climate Risk Management** 

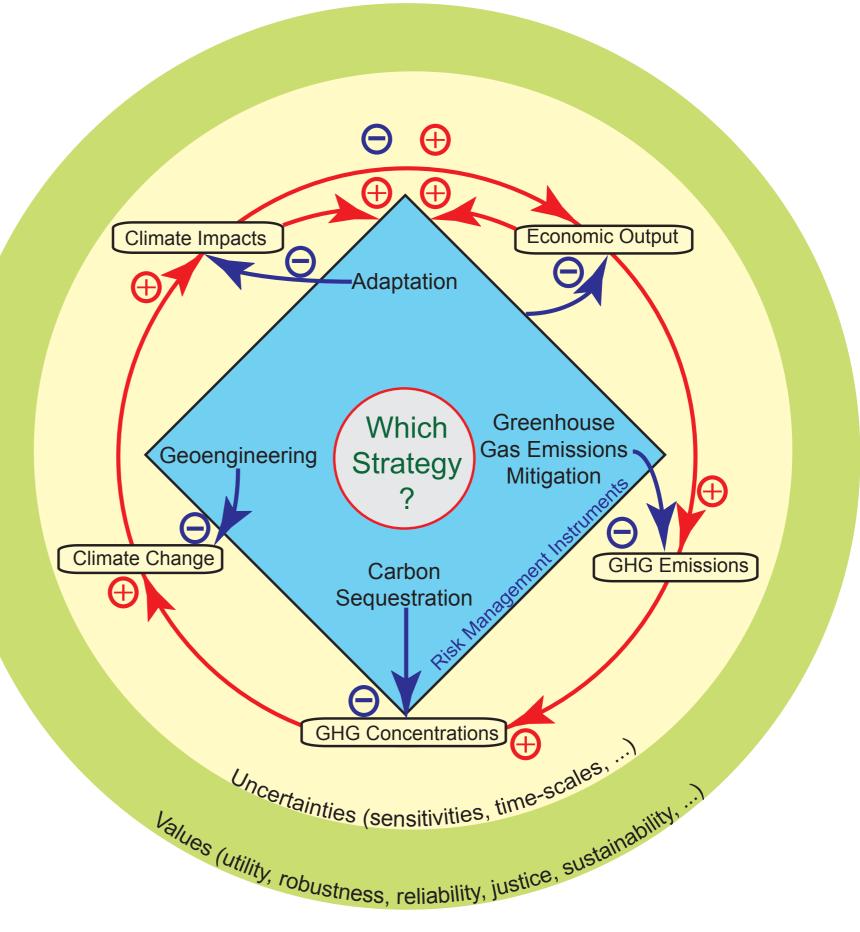
SCRIV

What are <u>sustainable</u>, <u>scientifically sound</u>, <u>technologically feasible</u>, <u>economically efficient</u>, and <u>ethically defensible</u> strategies for managing the risks associated with climate change?



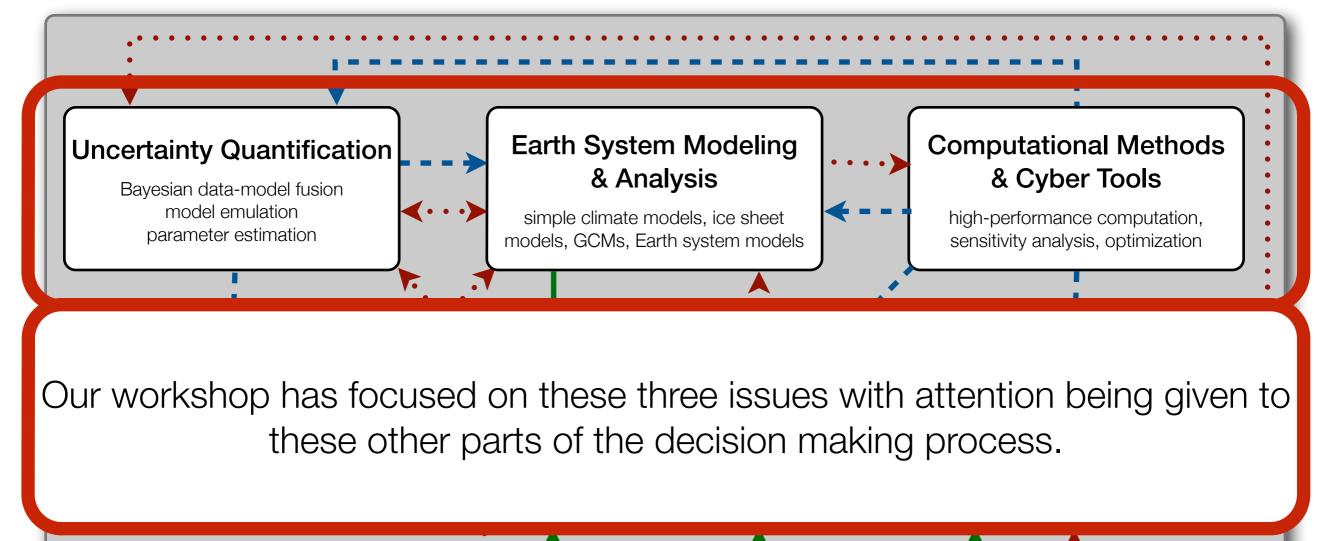
http://scrimhub.org

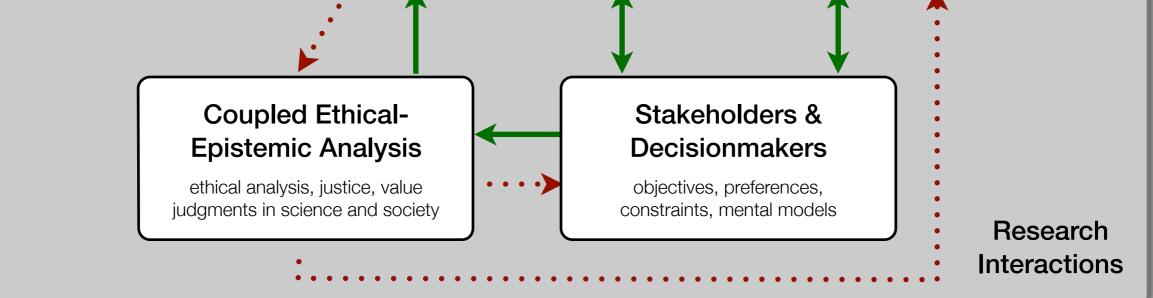
Assessing climate risk management strategies requires the analysis of complex systems with nontrivial interactions and emerging properties.

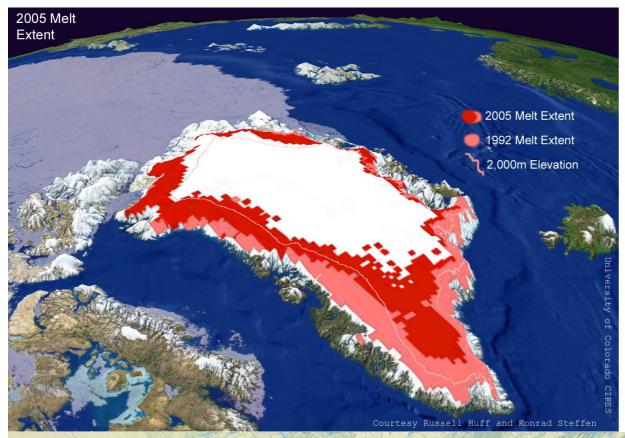


What are challenges?







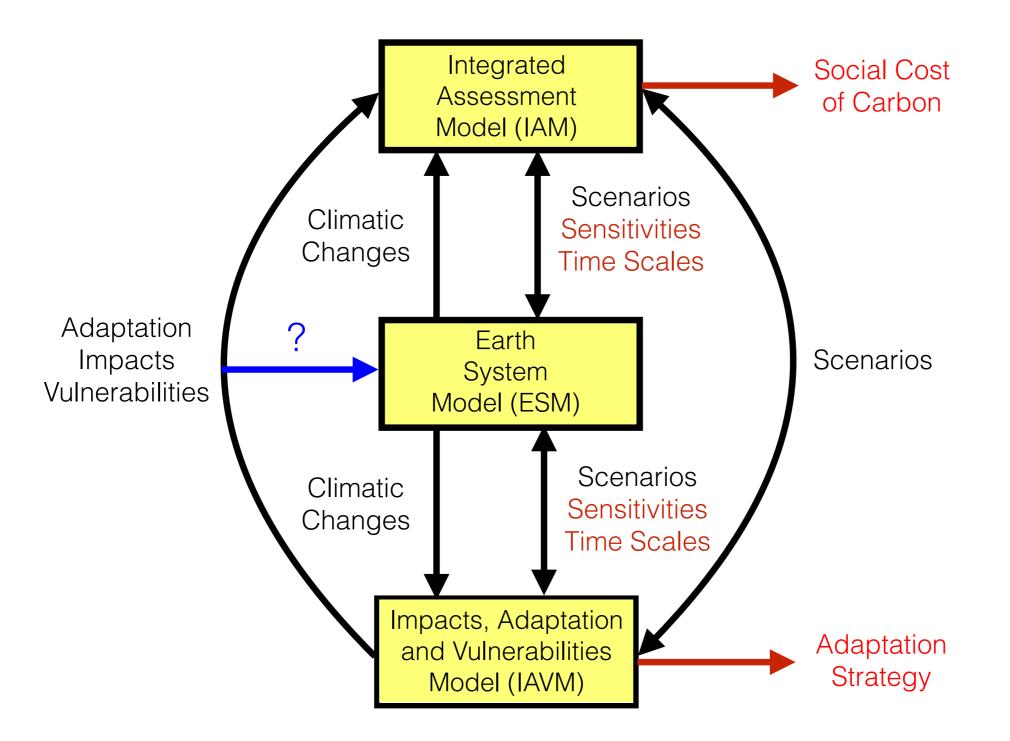


The Greenland Ice Sheet (GIS) is melting. A GIS disintegration could eventually raise the global mean sea level by roughly seven meters.

#### So what?

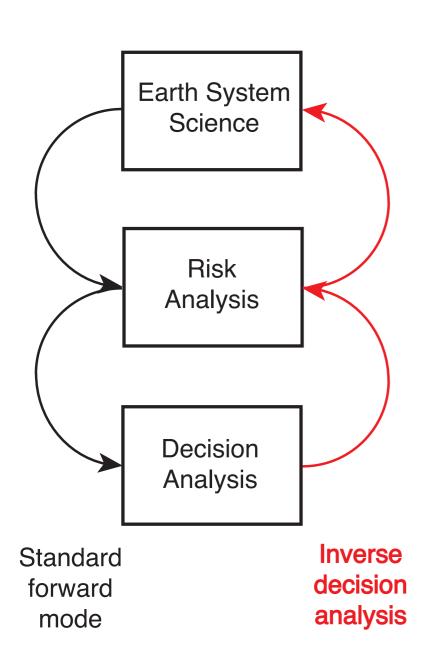


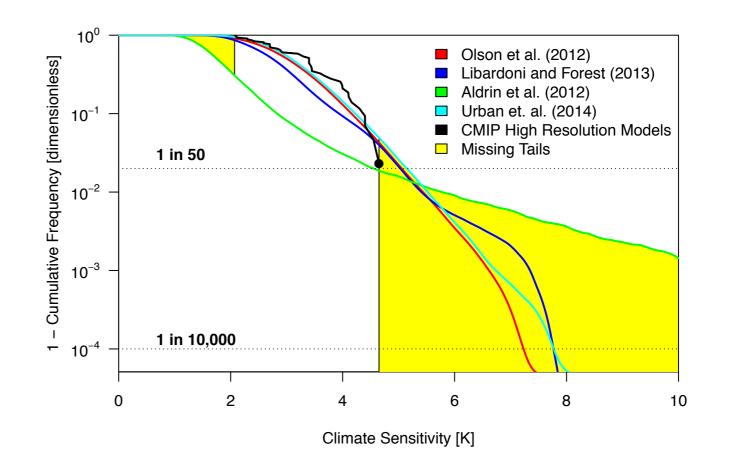
Analyzing climate risk management strategies requires a **tight and bidirectional linking** of disciplines and models.



6<sup>8</sup>

### Inverse decision analysis can point to promising and decision-relevant research directions





- Which uncertainties matter the most for decision making?
- What constrains the fat tail of climate sensitivity?
- How fast can the Greenland Ice Sheet disintegrate?

## This brings us back to some of the tools and techniques being discussed here...

## Fundamental uncertainties exist for projections of future climate

The uncertainties fall into a few broad categories:

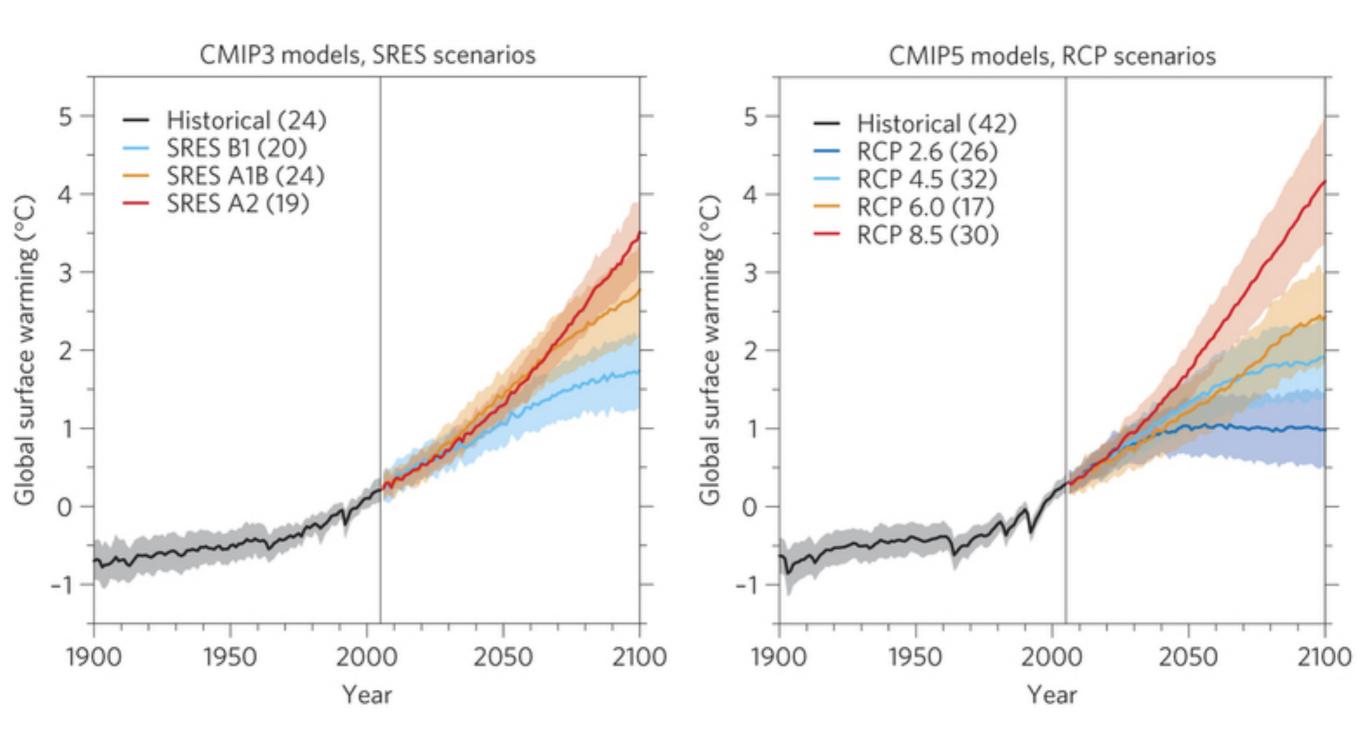
- observational uncertainty
- forcing or scenario uncertainty
- model uncertainty  $\Rightarrow$  structural, parametric (i.e., right physics, right settings)
- natural/internal variability

#### Goals

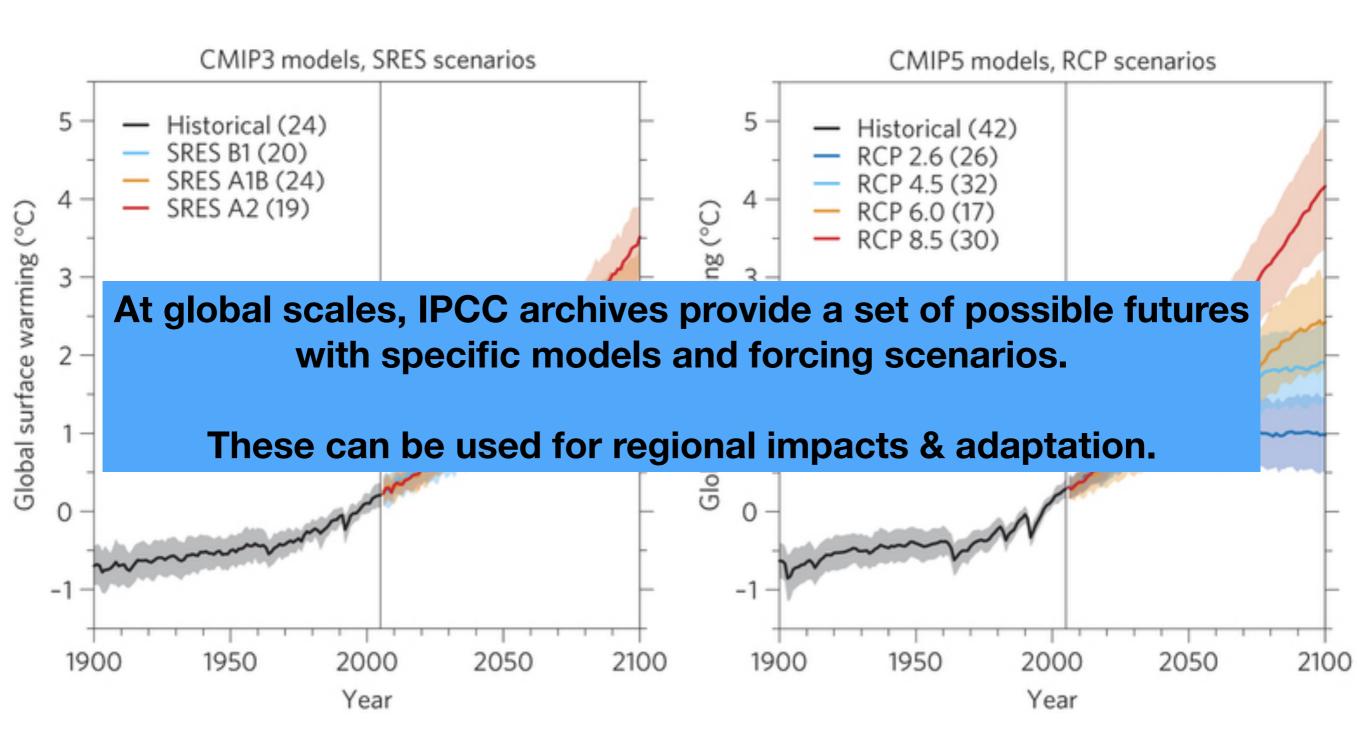
1. Can we separate uncertainties between Global v. Regional response?

- 2. How do we compare ensemble approaches? MME v. PPE v. ICE
- 3. How does structural uncertainty in regional changes assessed?

#### global mean temperature projections in CMIP3 and CMIP5

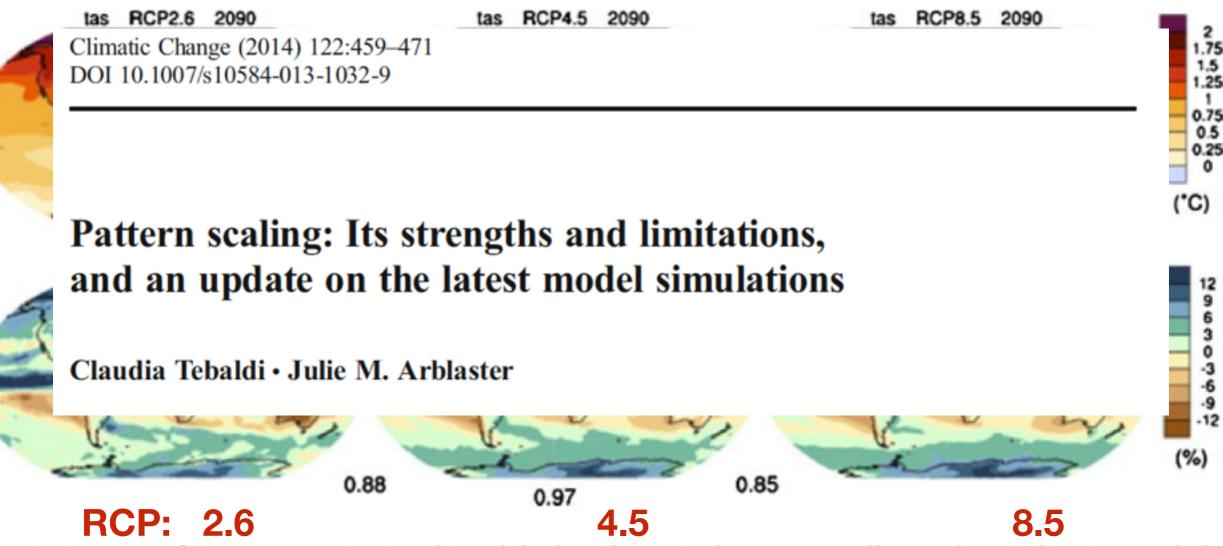


#### global mean temperature projections in CMIP3 and CMIP5



### The Pattern Scaling approach: A Method for Regional Model Predictions

P(x, y, z, t) = T(t)p(x, y, z)



**Fig. 1** Patterns of Temperature (*top*) and Precipitation (*bottom*) changes according to three RCPs (we exclude 6.0 because only a small number of models ran it). The numbers between adjacent maps show the pattern correlation between the two maps. The *lower number* in the center shows the pattern correlation between the first and third map

#### Tebaldi & Arblaster (2014, Climatic Change)

## The Problem: Model Predictions have multiple sources of uncertainty...

 Epistemic uncertainty: getting the model right (the right model physics & right structure)

Example Work at Penn State: Developing Global Teleconnection Operators (GTOs) to estimate structural uncertainty in regional response to global Sea Surface Temperature patterns

Li, Forest, & Barsugli (2012, J. Geophys. Res.) Comparing two methods to estimate the sensitivity of regional climate simulations to tropical SST anomalies

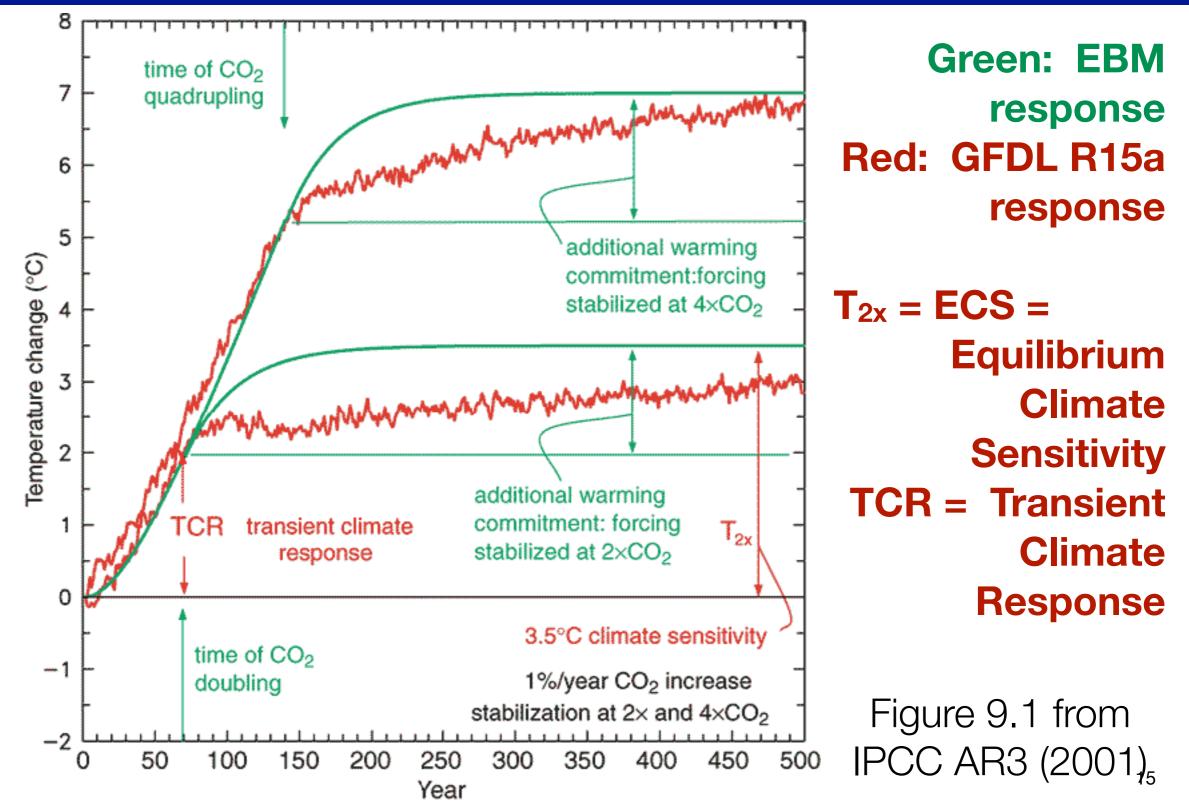
**Tsai, Forest, & Wagener (2014, Clim. Dynamics)** Estimating the regional climate responses over river basins to changes in tropical sea surface temperature patterns

# What matters for long-term climate prediction?

- Controls on:
  - Long-term warming
  - Delay by ocean
  - Net forcing

- Uncertainties in:
  - Climate Sensitivity
  - Rate of Ocean Heat Uptake
  - Forcing by: Aerosols, Carbon-cycle, Landuse, Natural GHG Emissions, etc.

## Climate Sensitivity is the equilibrium response of the climate to a constant forcing



### **Major Climate Projection Uncertainties**

Consider the energy balance equation for the change in global-mean surface temperature ( $\Delta T$ ) from equilibrium:

$$c_p \frac{d\Delta T(t)}{dt} = F(t) - \lambda \Delta T(t) - \Phi_o(K_v)$$

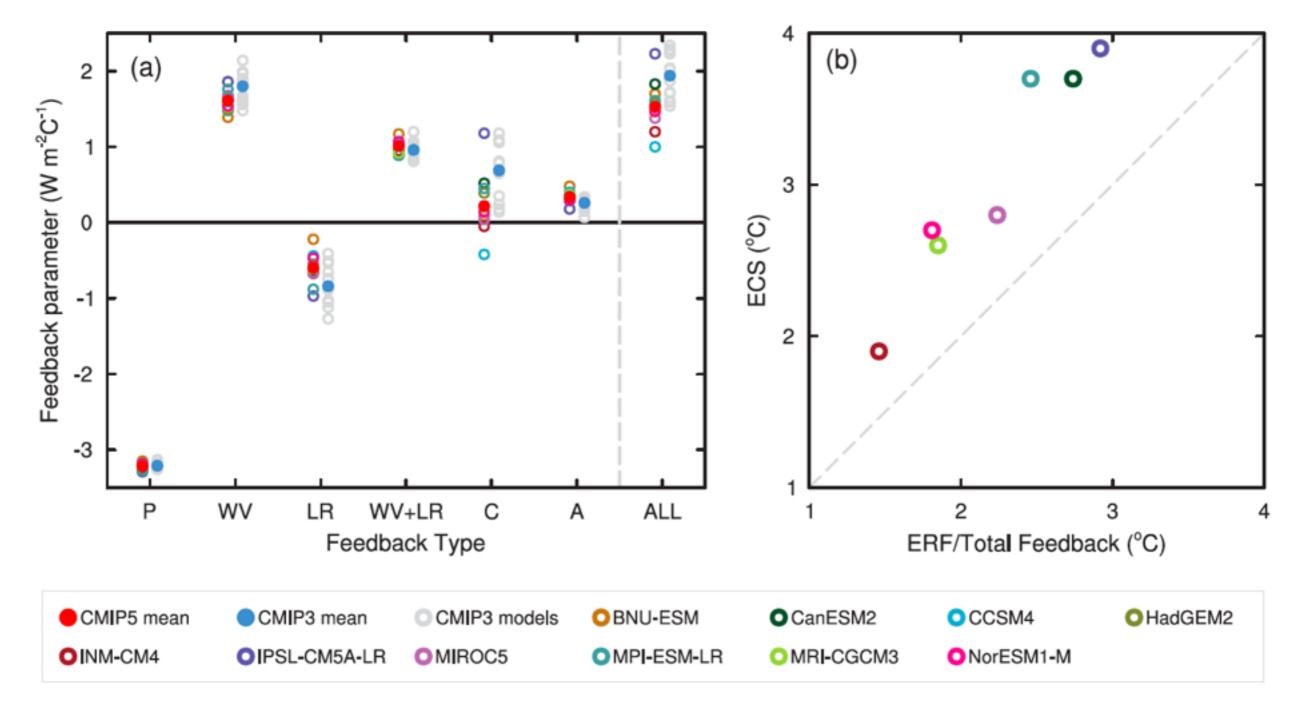
Change in globalFutureNetFlux of heatmean heat contentForcingsFeedbacksinto deep- $\lambda = 1/S$ ocean

**Conceptually:** This is a good framework for organizing where the uncertainty exists.

In practice: For state-of-the-art models, each uncertainty is an aggregate quantity and cannot be identified with any one specific model component or process.

### Uncertainty in Atmospheric Model Feedbacks (IPCC WG1 AR5 Figure 9.43)

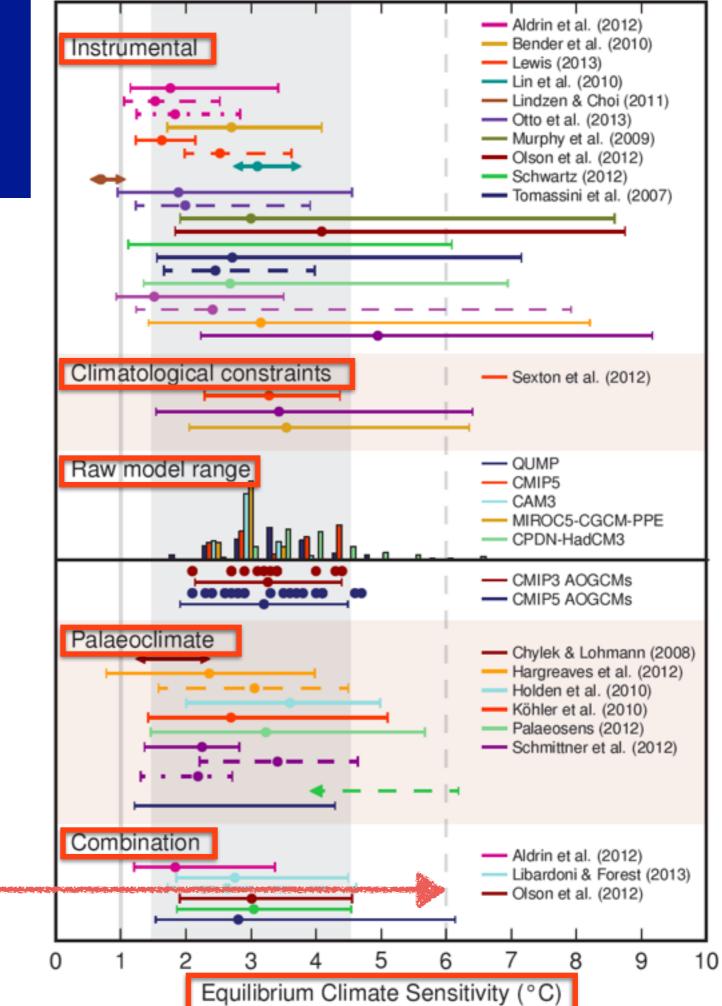
 Uncertainty in Planck, Water Vapor, Lapse Rate, Cloud, Albedo, and ALL Combined

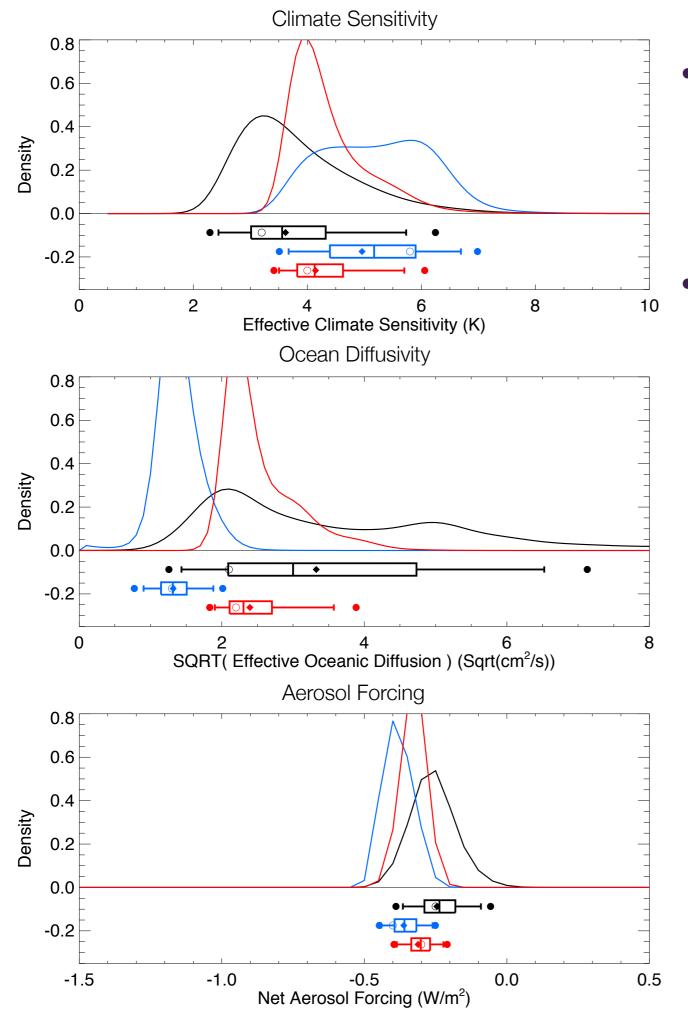


### Uncertainty in Equilibrium Climate Sensitivity

- IPCC Range: 1.5-4.5 °C
- Box 12.2 Figure 1
- Multiple approaches based on different methodologies and observational data/ proxies.

Work by our groups:





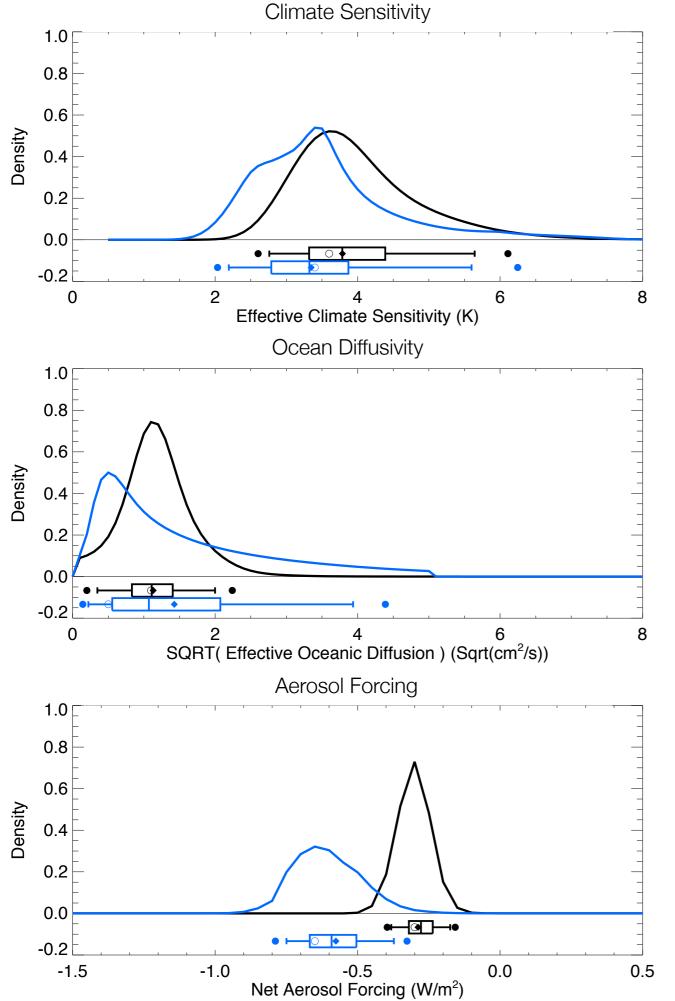
## How are the marginal PDFs impacted?

 Extending model diagnostics with the new model

All Diagnostics start in 1941

Diagnostics end in 1990 Diagnostics end in 2000 Diagnostics end in 2010

Libardoni, Forest, & Sokolov (In prep)



## How are the marginal PDFs impacted?

Changing the model and forcings

New Model (MESM) Old Model (IGSM) (Libardoni & Forest, 2013)

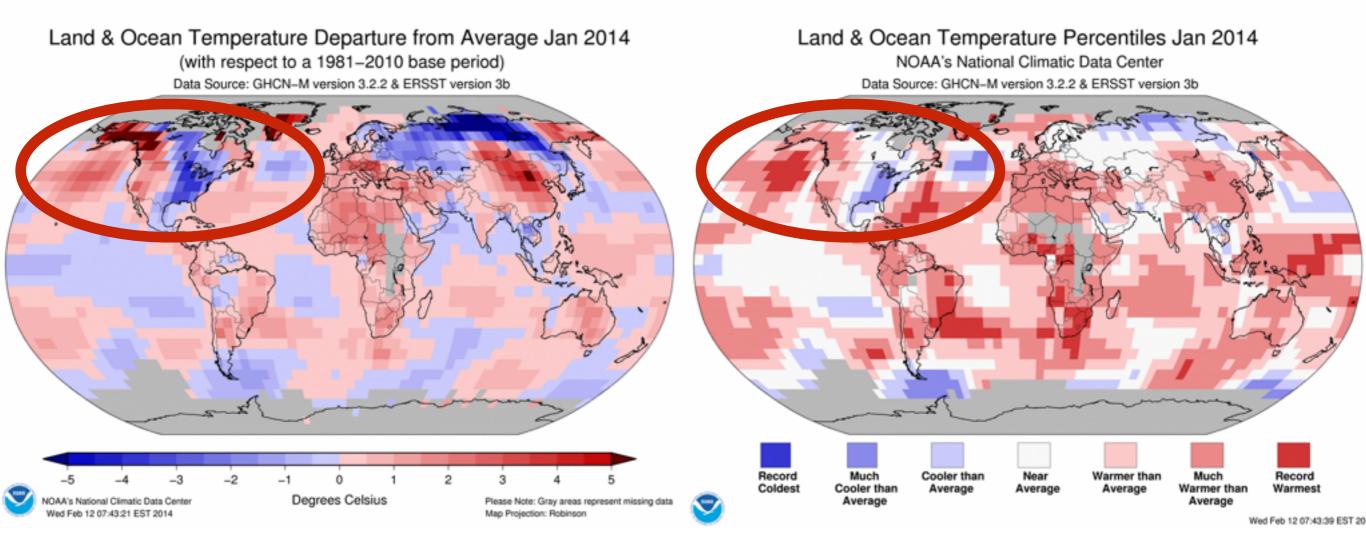
> Both use Same Diagnostics

Libardoni, Forest, & Sokolov (In prep)

# So this begins to address global scale uncertainties...

How do we start assessing uncertainties at regional scales?

#### What are Drivers of Regional Extremes? Example: Polar Vortex in January 2014



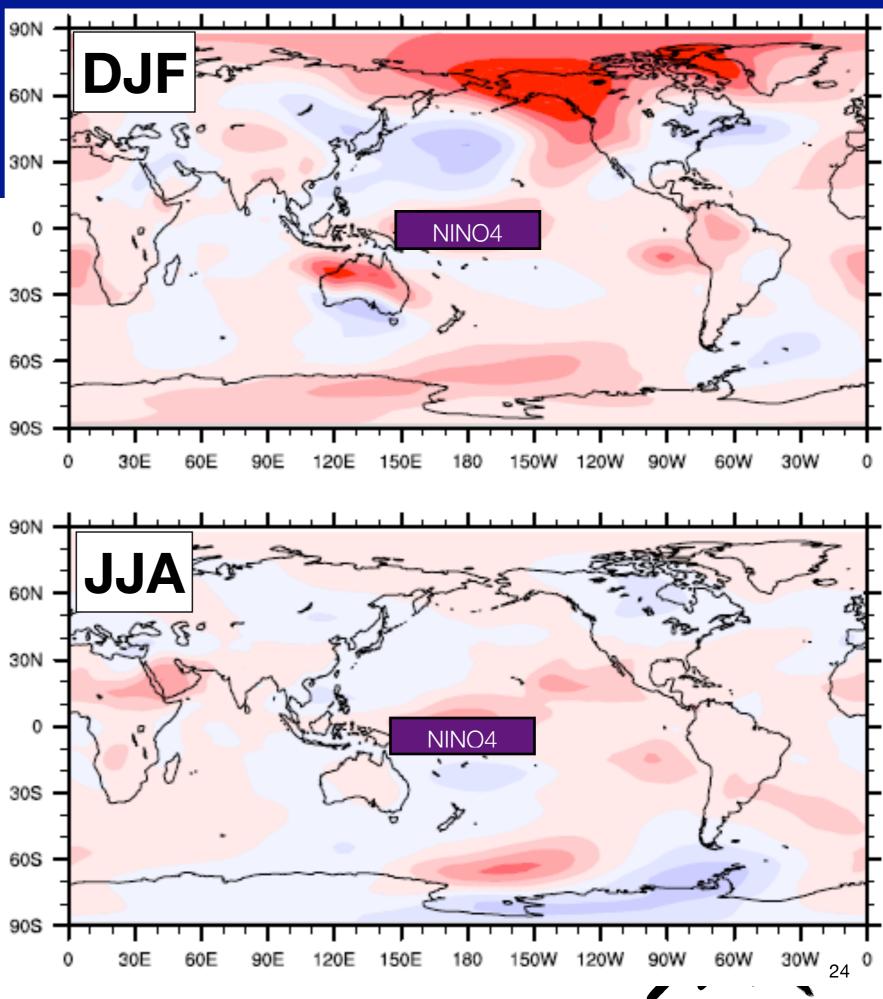
SST patterns are a primary driver with land surface and sea-ice potentially being as important. Implies that regional changes require better estimates of ocean variability.

## Can we quantify structural uncertainty in model response at *regional scales*?

- Alternatively, how can we go beyond the IPCC/CMIP multi-model ensemble to assess climate risk?
- How: Create *idealized* experiments with "known" forcings to provide metrics (or framework) for comparing model response <u>at regional scales</u>
- Purpose: Estimate the teleconnection response that adds to the mean climate response pattern.
  - *NB:* According to the physics community, teleconnections are second-order cumulant stats.

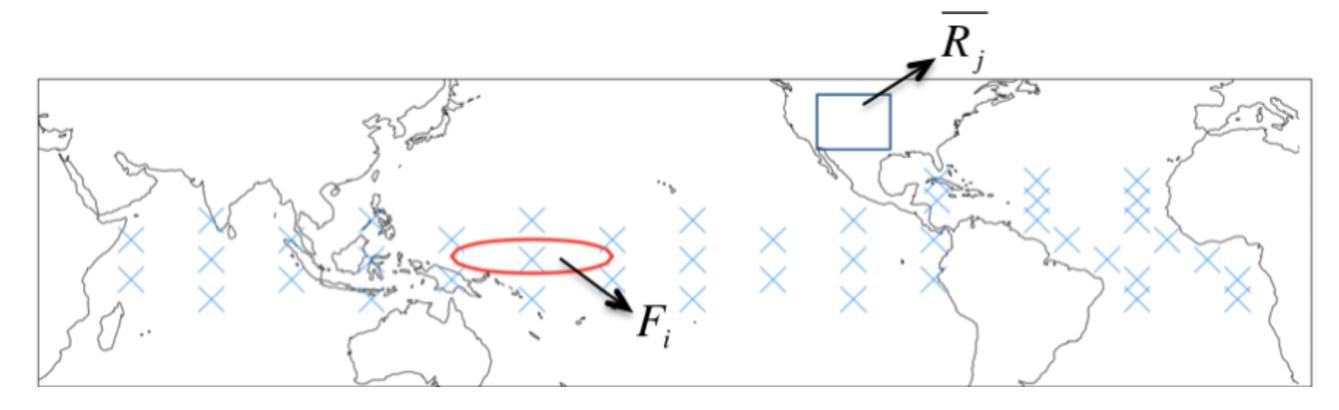
### Example: Teleconnection

- Response to Nino4
- Seasonal Mean
   Temperature on
   850hPa surface





### **Global Teleconnection Operator: GTO** Estimate the ensemble-mean response, $R_j$ , to the $\Delta$ SST forcing, $F_i$



Estimate Global Teleconnection Operator, K<sub>ii</sub>, from:

$$\overline{R_j} = K_{ij} \cdot F_i + arepsilon$$

Repeat this estimate for all SST anomaly locations in tropics **Patch Method:** Barsugli and Sardeshmukh (2002, J. Climate) **Random Patch Method:** Li, Forest, & Barsugli (2012, JGR-A) <sub>25</sub>

### GTO Experiments (X = complete, O = in progress)

	Models						
Resolution		NCAR	GFDL	HadAM3			
	CAM3.1	CAM3.5	CAM4	CAM5	AM2	(CPDN)	
T31				X			
T42	X						
T85	X						
FV0.9x1.25				X			
FV1.9x2.5	X	Χ	Χ	X	<b>X</b> *	<b>O</b> *	
FV4x5				X			
HOMME_N30				X			

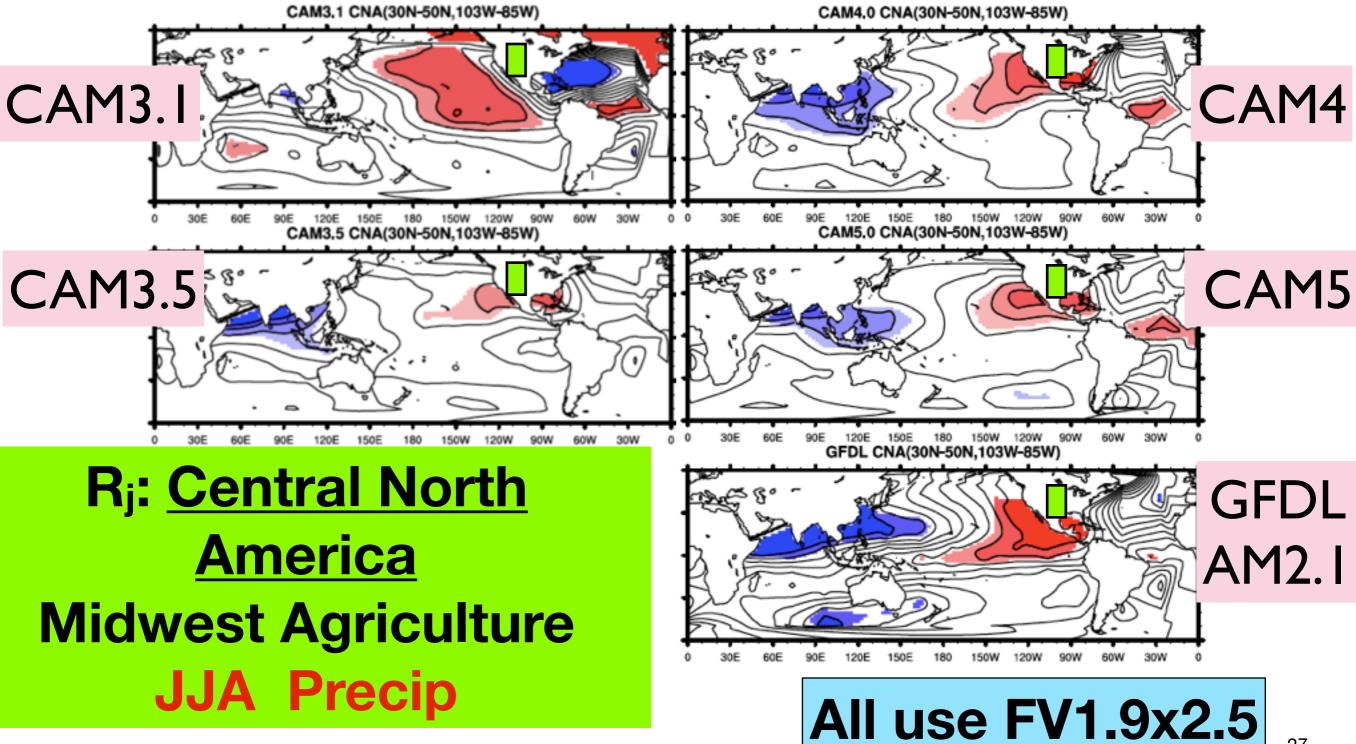
NOTE1: Ensemble size with n=400 is typically sufficient.

\*GFDL AM2 (2.0x2.5), HadAM3 (2.5x3.75)

NOTE2: GTO for HadAM3: a new ClimatePrediction.net experiment, >10k simulations

#### Testing for Structural Uncertainty GTO: Sensitivity Maps (*K*<sub>ij</sub>)

#### Comparing Physics



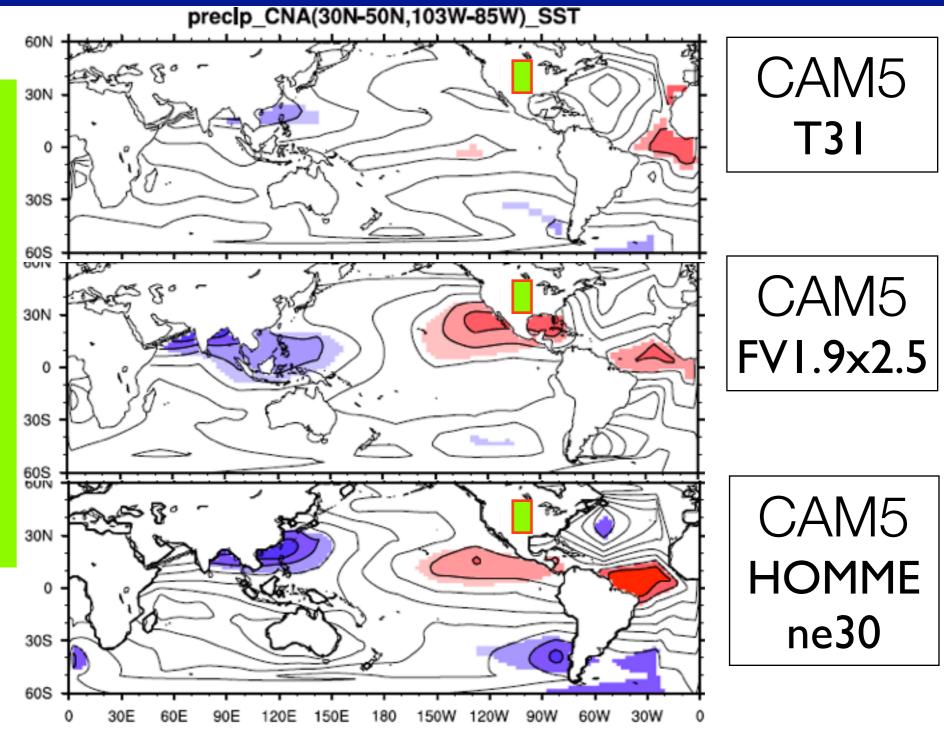
# Testing for Structural Uncertainty GTO: Sensitivity Maps (*K*<sub>ij</sub>)

#### Comparing DyCore

*R<sub>j</sub>*: Central North America

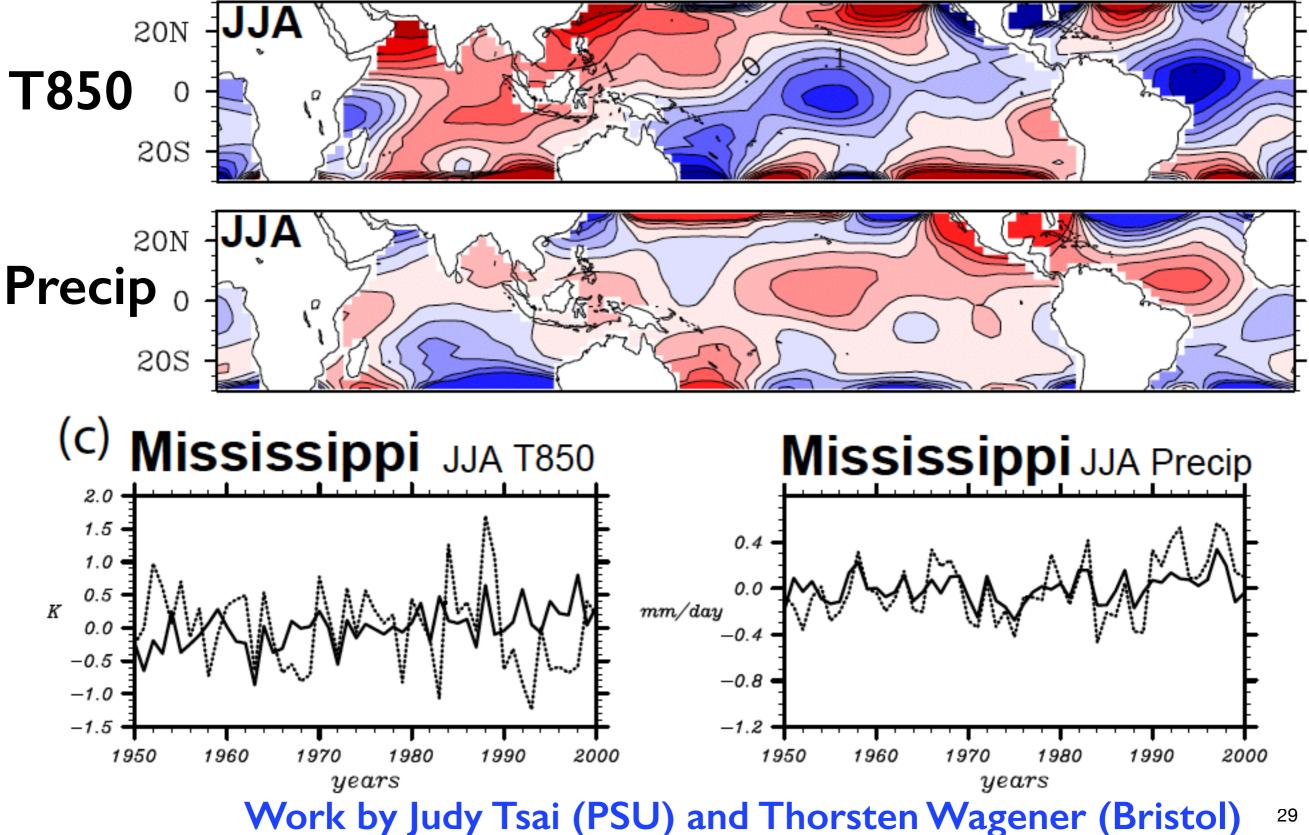
### Midwest Agriculture

**JJA Precip** 

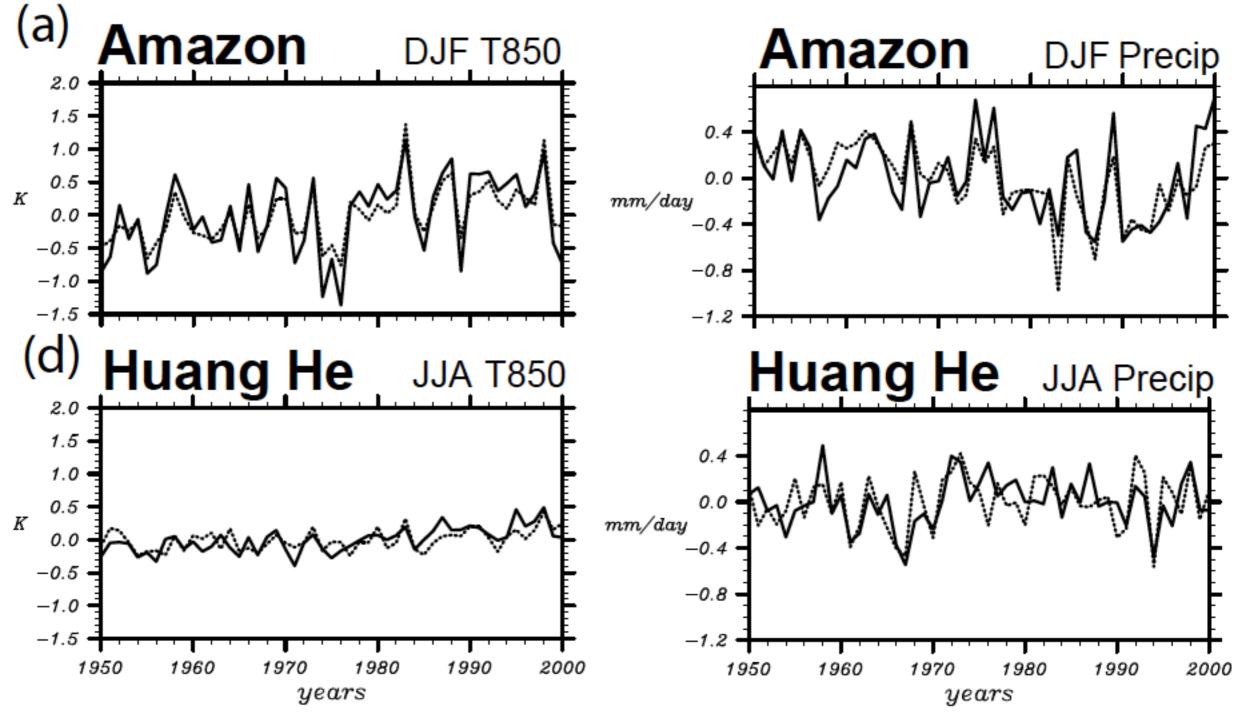


-22.4	-11.2	0	11.2	22.4

#### **Application to Mississippi River Basin (JJA)**



Application to Amazon River and Huang He Reconstruct T and Precipitation based on Green's function response to observed SST:  $R(t) = K_{ij}*SST(t,x_i)$ 

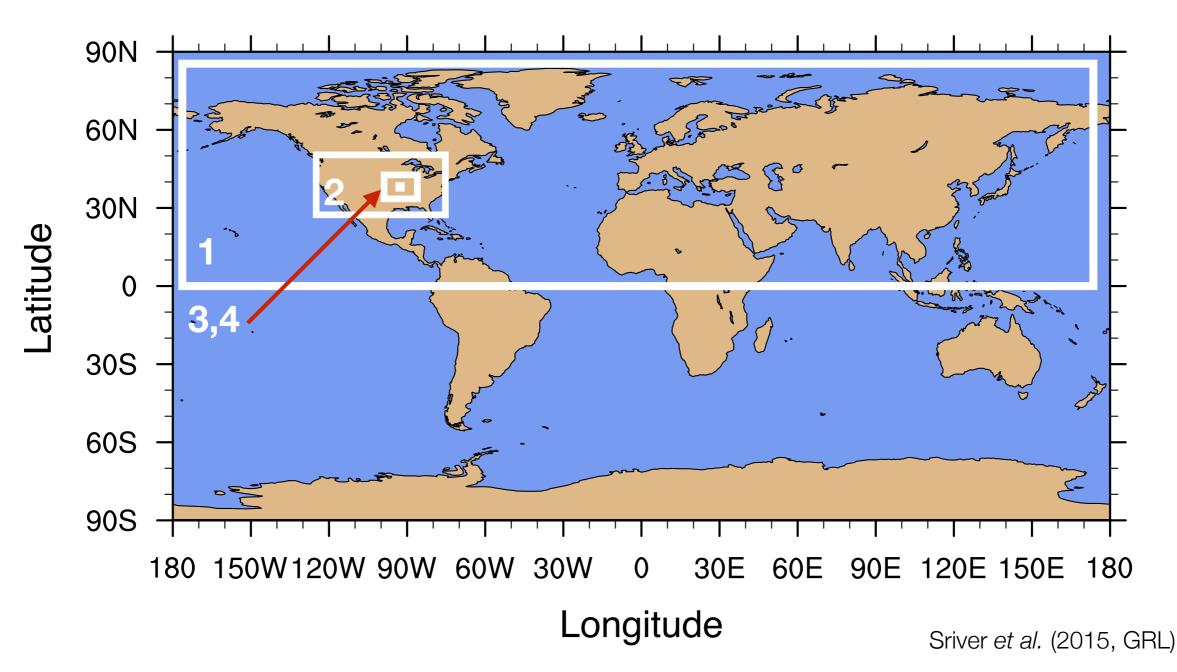


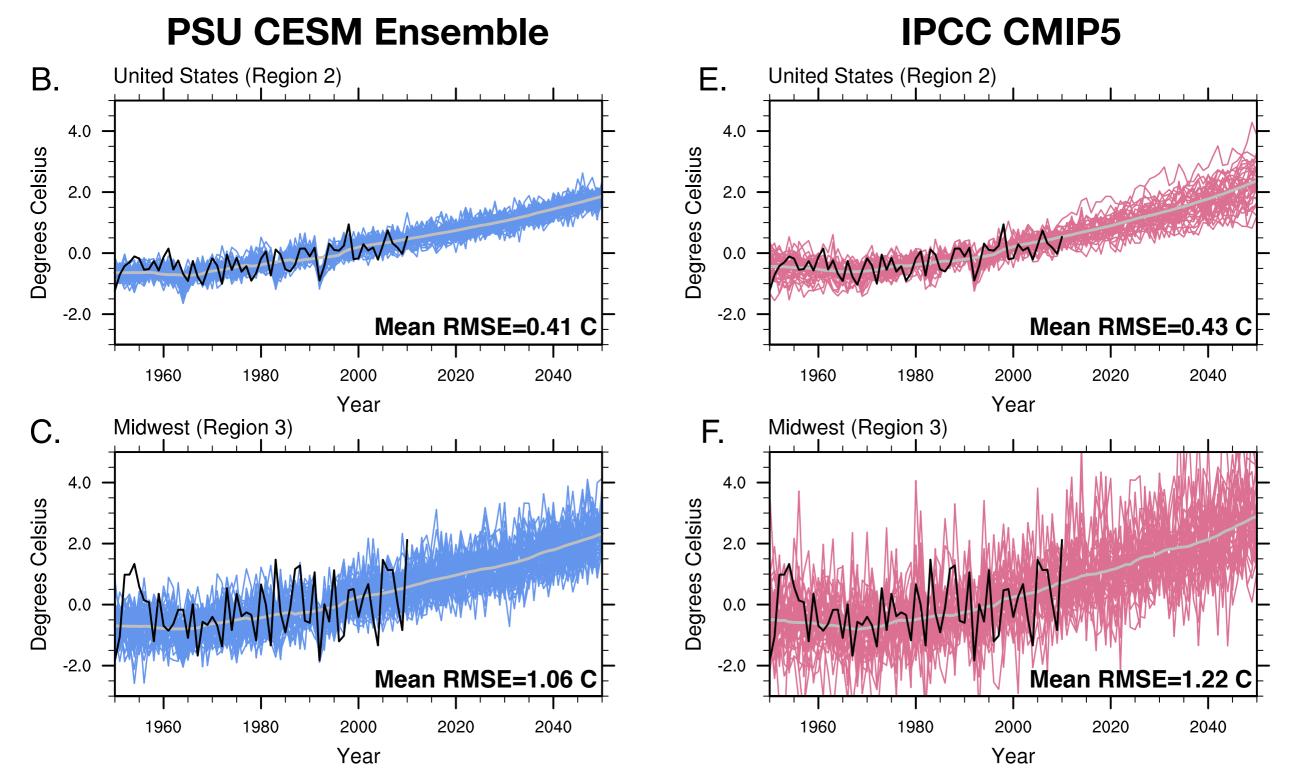
Work by Judy Tsai (PSU) and Thorsten Wagener (Bristol)

Goal: To add additional information on initial condition uncertainties

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#### Focus on representing climate information at regional scales





1.50 Member Initial Condition Ensemble: 1850-2100
2.Based on unique initial condition from 5000 yr control
3.Total Simulation years: 17,500 (150 h Srive

Sriver *et al.* (2015, GRL) 33

### there are fundamental uncertainties about projections of future climate

The uncertainties fall into a few broad categories:

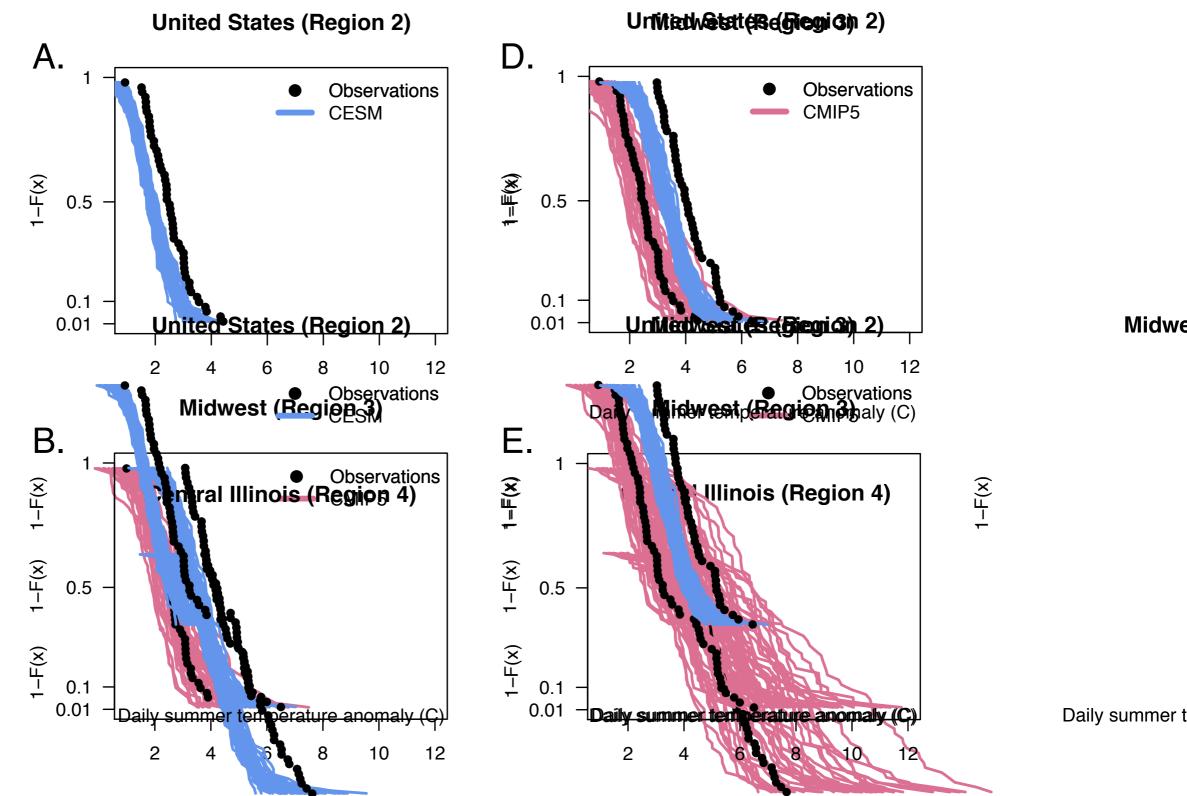
- observational uncertainty
- forcing or scenario uncertainty
- model uncertainty  $\Rightarrow$  structural, parametric
- natural/internal variability
- The magnitudes of the projection uncertainties vary in space and time; smaller spatial and temporal scales  $\Rightarrow$  greater uncertainties.

Climate projections may also exhibit systematic biases.

Failing to account for biases and the full range of uncertainties can lead to overconfident (and unhelpful) projections.

We are working to incorporate all these into decision-making frameworks

Next: How do we assess uncertainty at regional scales directly?



Survival Function for Summer block maxima of daily surface temperature anomalies (1956-2005), for the CESM and CMIPS in the cest of the ce

35

#### Summary

#### Goals

#### 1. Can we separate uncertainties between Global v. Regional response?

- Pattern scaling approaches suggest yes but formal assessment is needed.

#### 2. How do we compare ensemble approaches? MME v. PPE v. ICE

- Multiple estimates with similar results. Does each have a place?

#### 3. How does structural uncertainty in regional changes assessed?

- Proposing metrics for this uncertainty is difficult.
- Teleconnections (aka second-order cumulant statistics) are one option.

 The GTO estimates the first-order linear response to SST patterns and can be used for emulations.

#### Thank you! mailto:ceforest@psu.edu

### Questions?