

Joint ICTP-EAIFR-IUGG Workshop on Computational Geodynamics: Towards Building a New Expertise Across Africa



tectonic forces.
richard katz
university of oxford

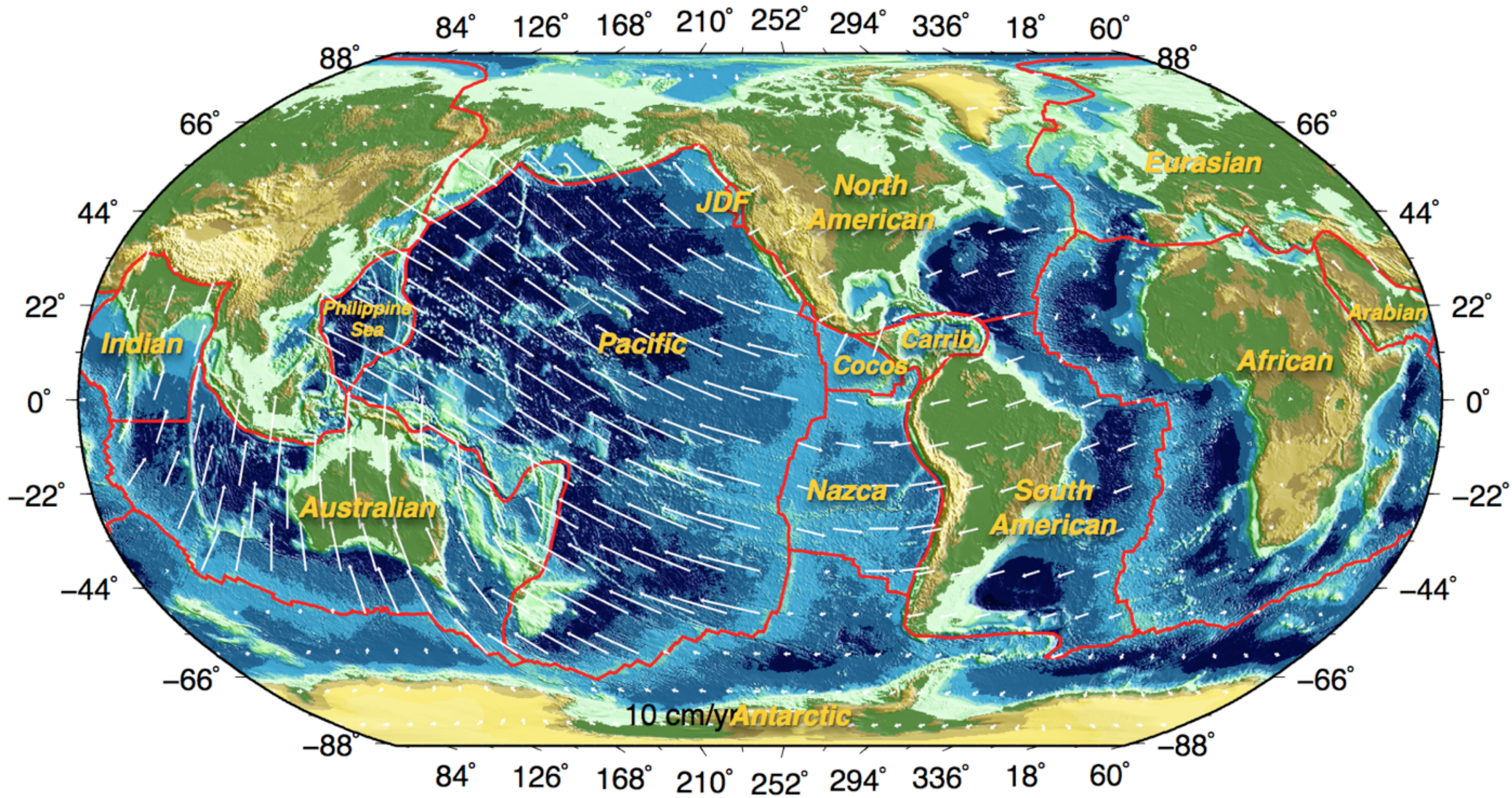
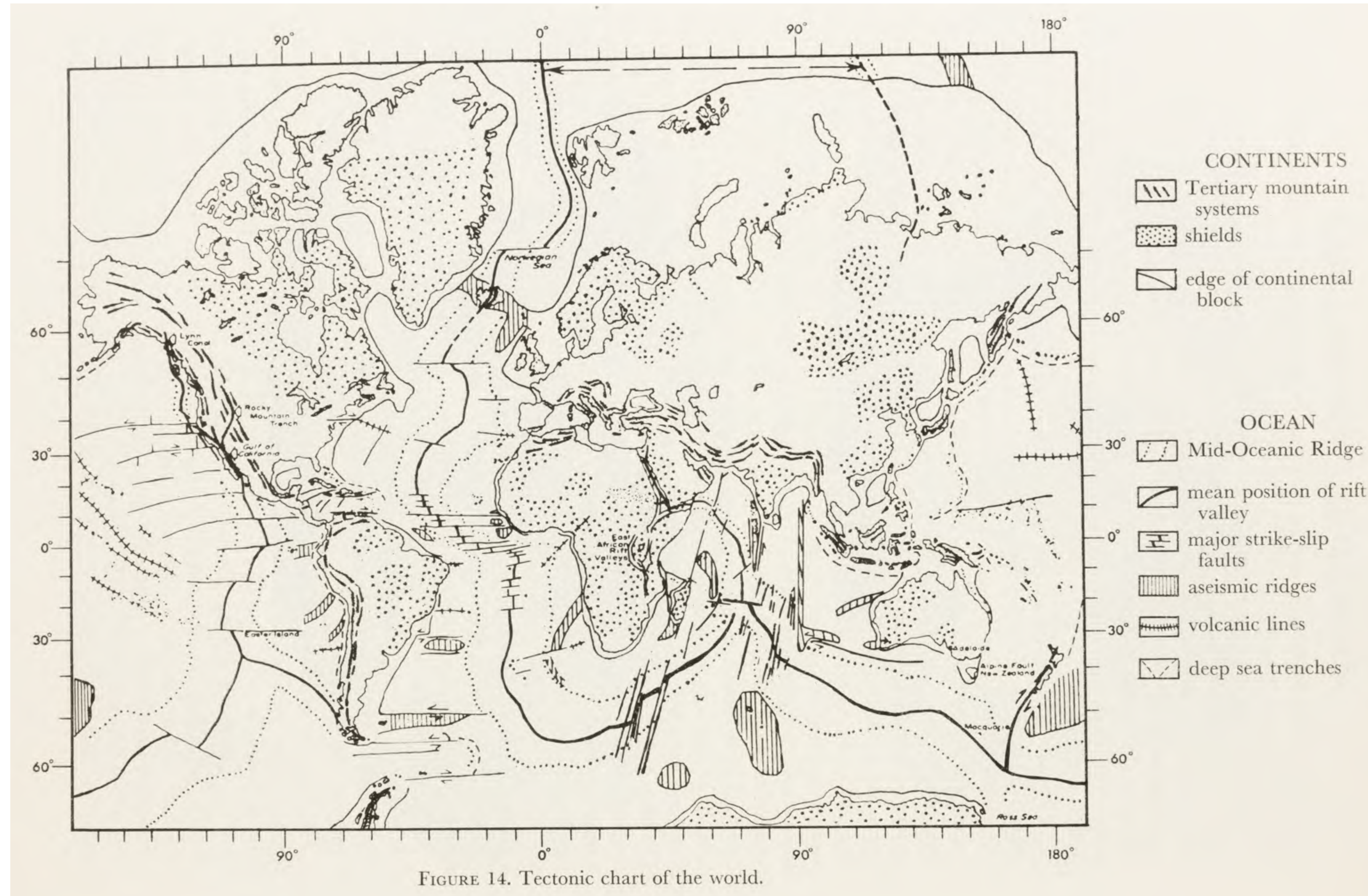


Plate tectonic theory in 1965



How big is Africa, really?



Plate tectonics as a *kinematic* theory

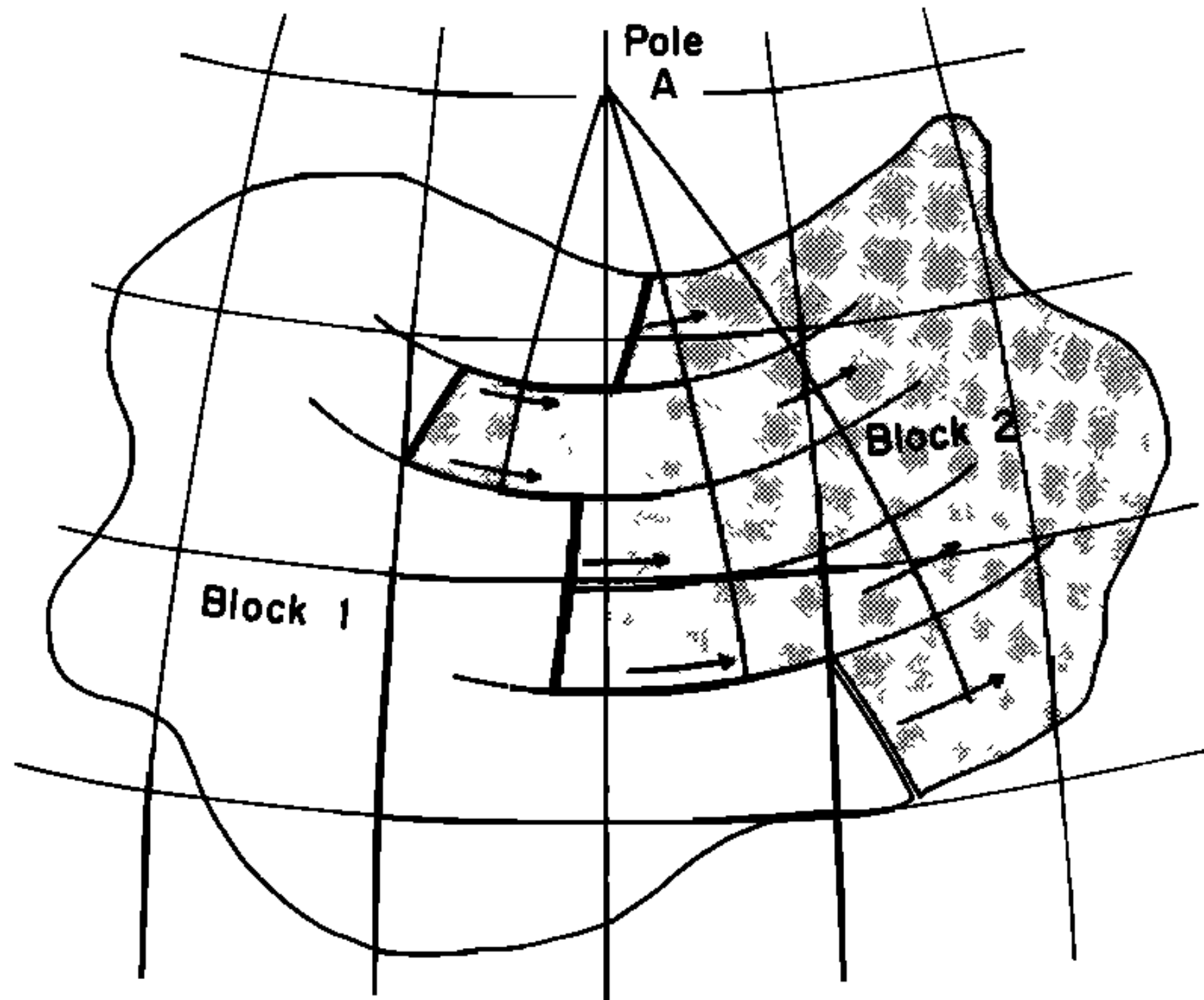
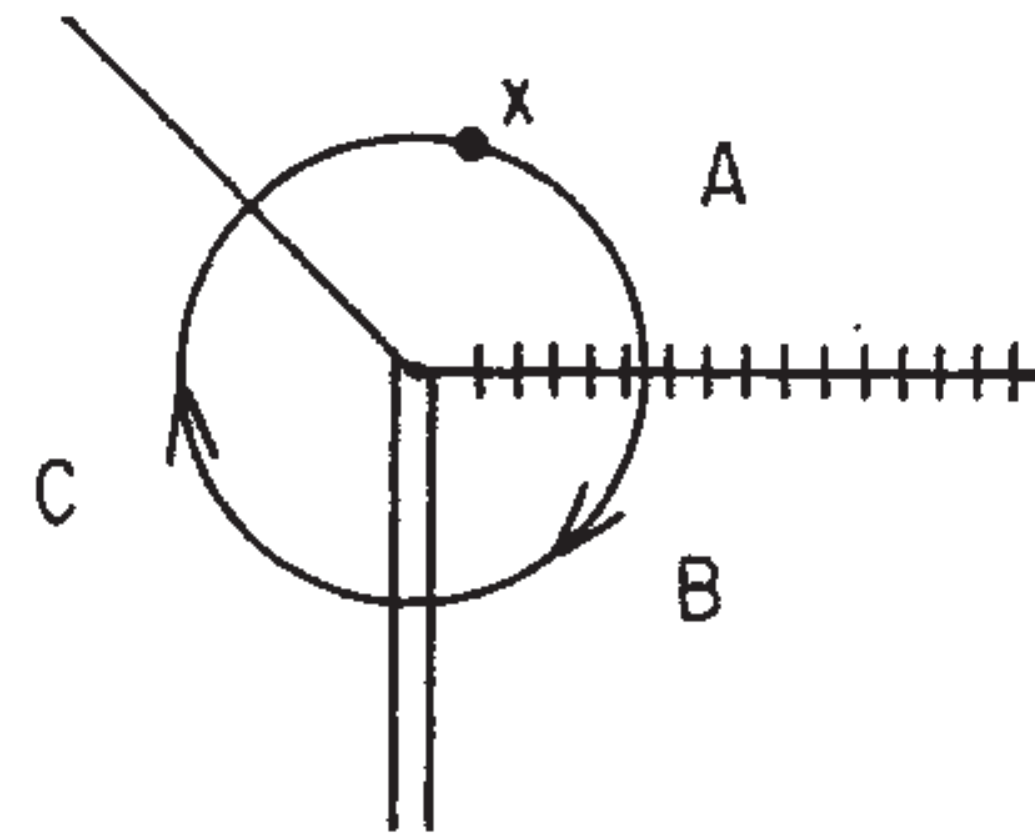


Fig. 4. On a sphere, the motion of block 2 relative to block 1 must be a rotation about some pole. All faults on the boundary between 1 and 2 must be small circles concentric about the pole A.



- ==== RIDGE
- + + + + + TRENCH
- TRANSFORM FAULT

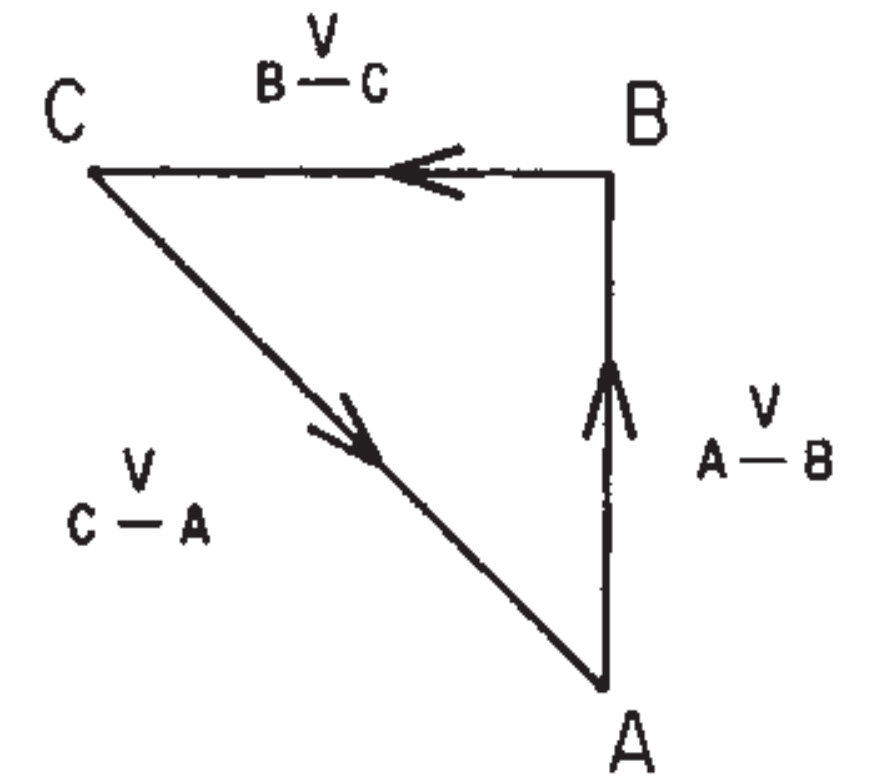
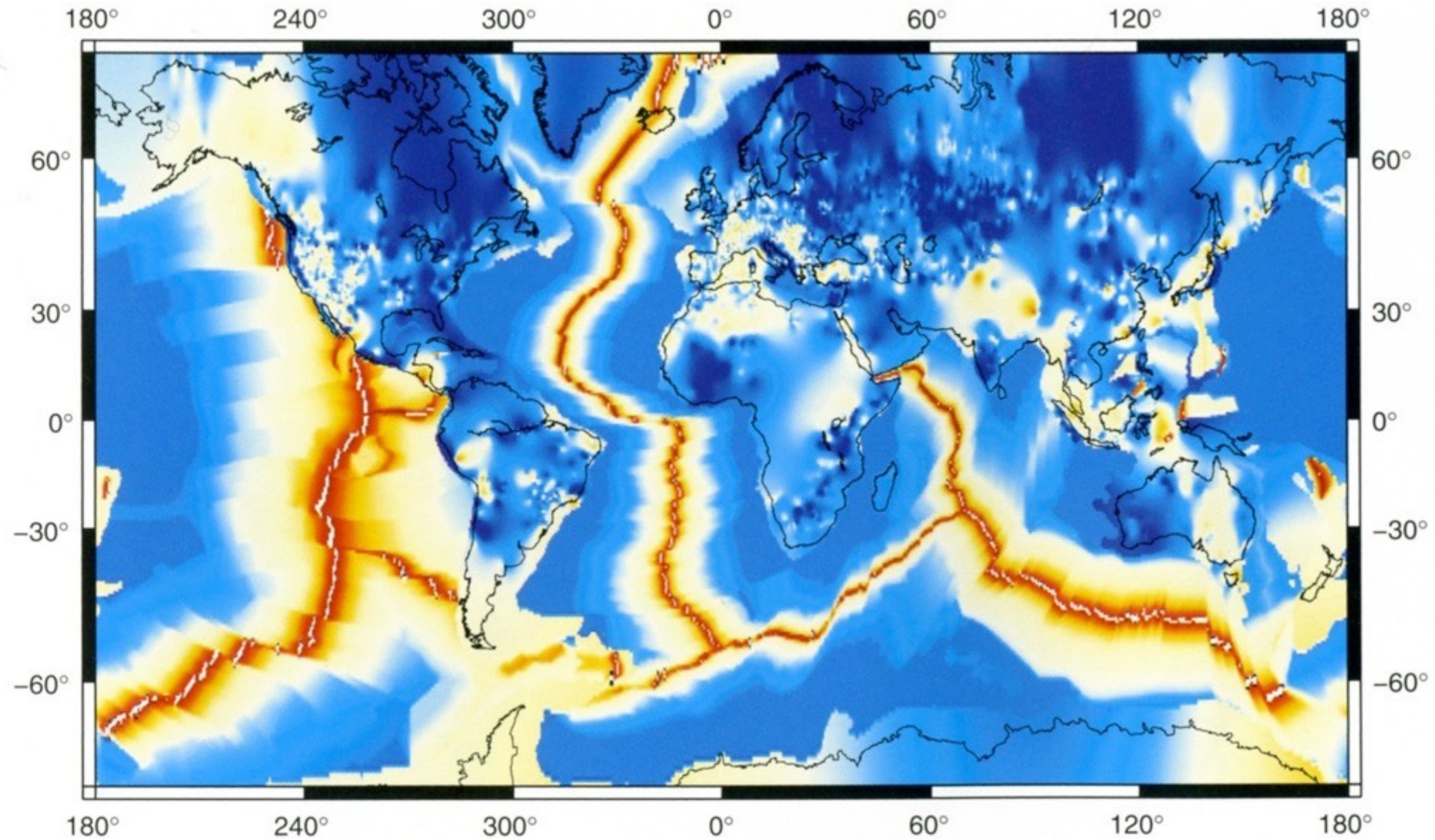
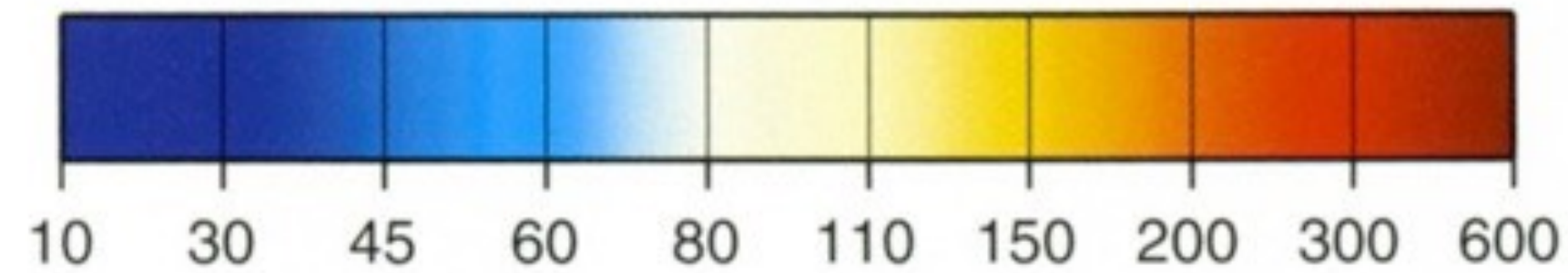


Fig. 1. The circuit and its vector diagram show how a ridge and a trench can meet to form a transform fault.

Plate tectonics controls heat flow



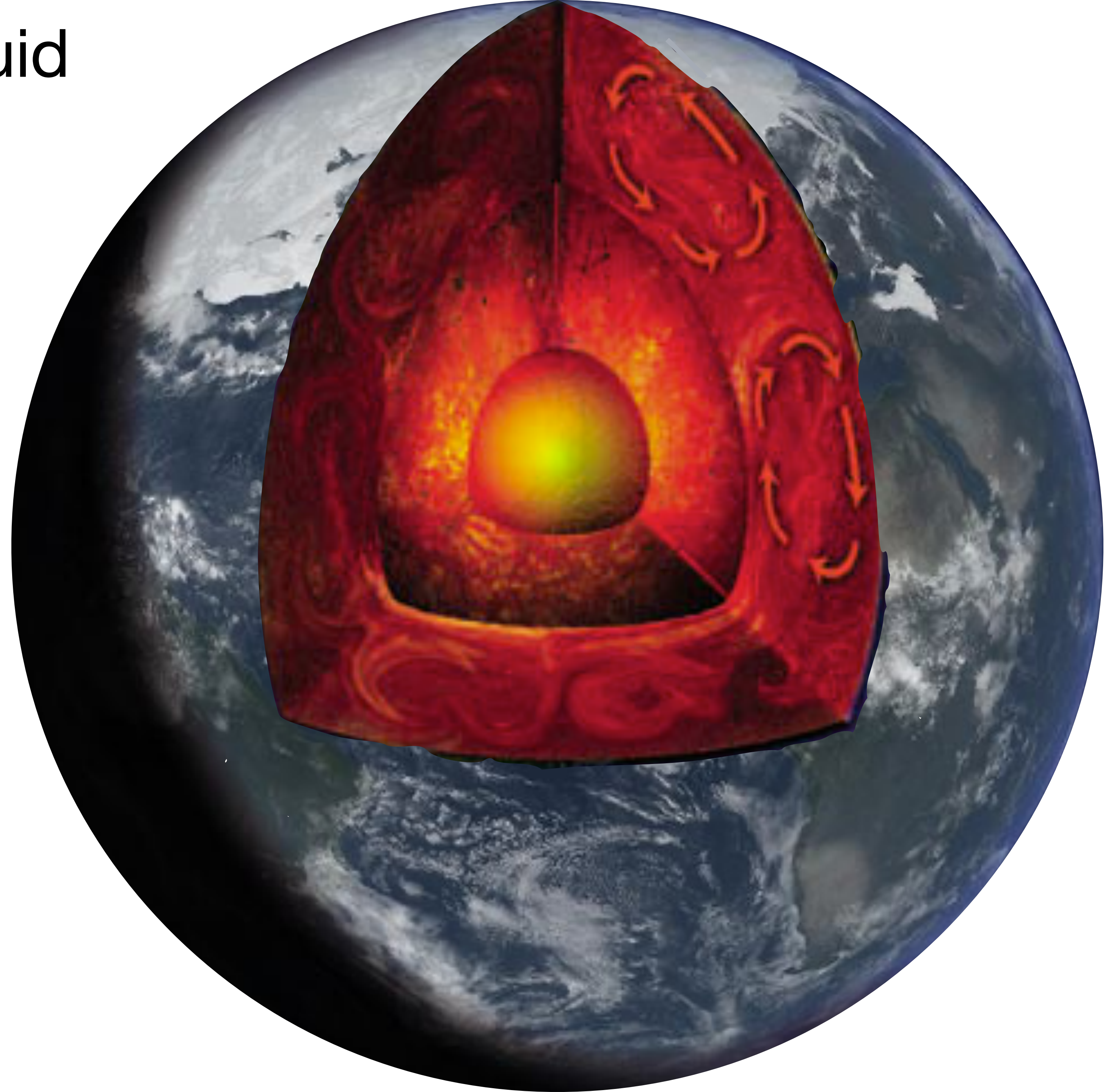
Jaupart &
Mareschal



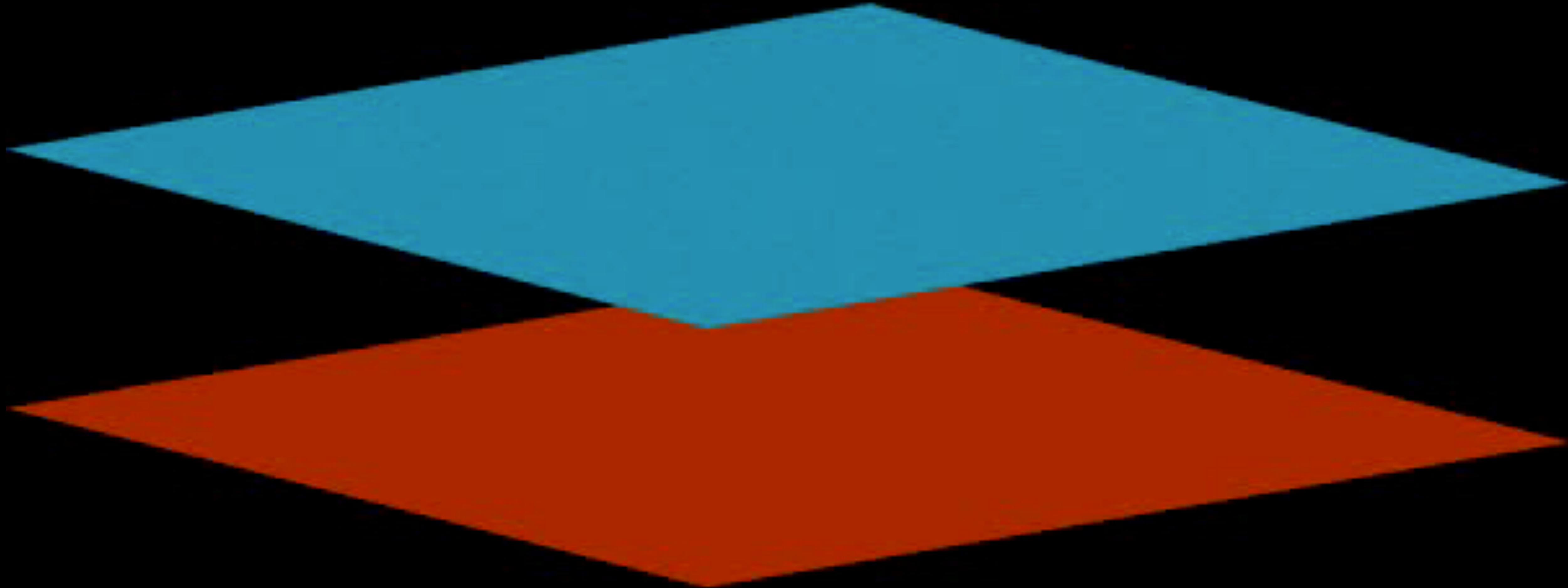
mW/m²

The solid mantle is a fluid
(on long timescales)

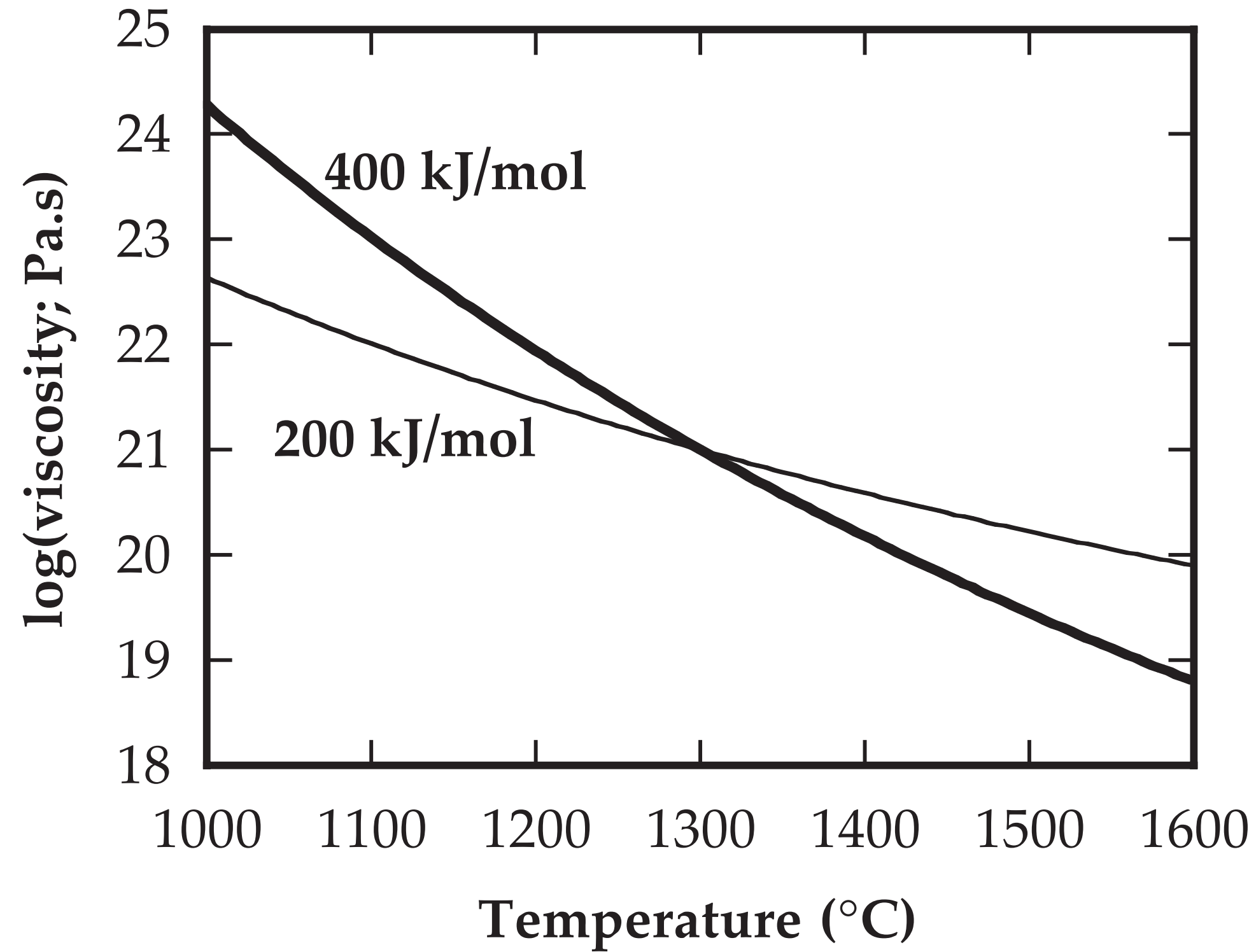
(Solid ice is a fluid too)



Isoviscous thermal convection far above critical Ra



Mantle viscosity, plates and convection



$$\eta(T) = \eta_0 \exp \left[\frac{E^*}{R} \left(\frac{1}{T} - \frac{1}{T_0} \right) \right]$$

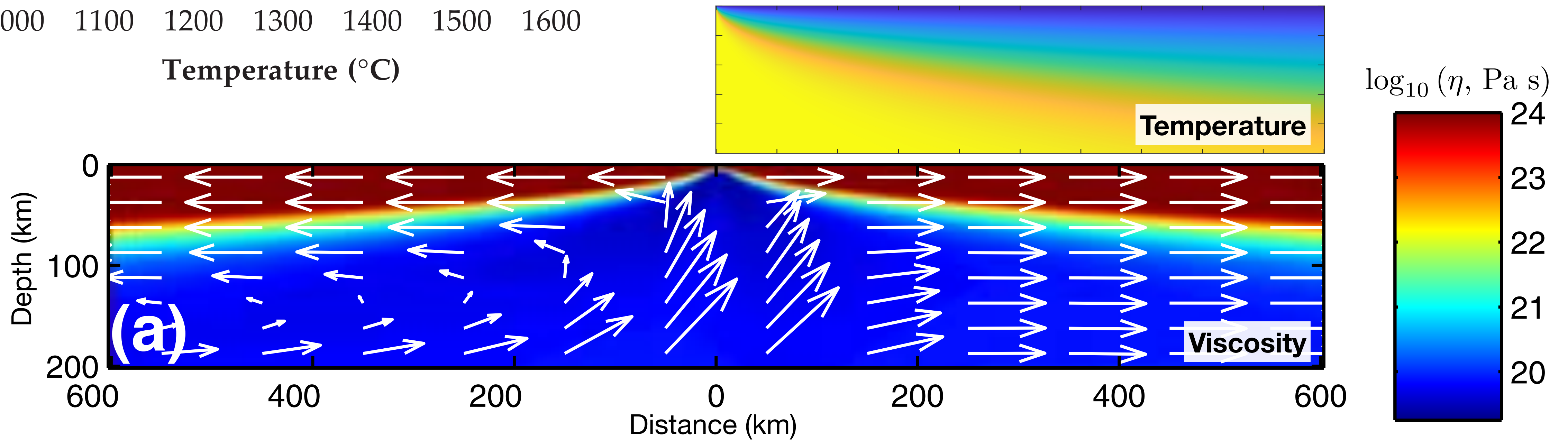
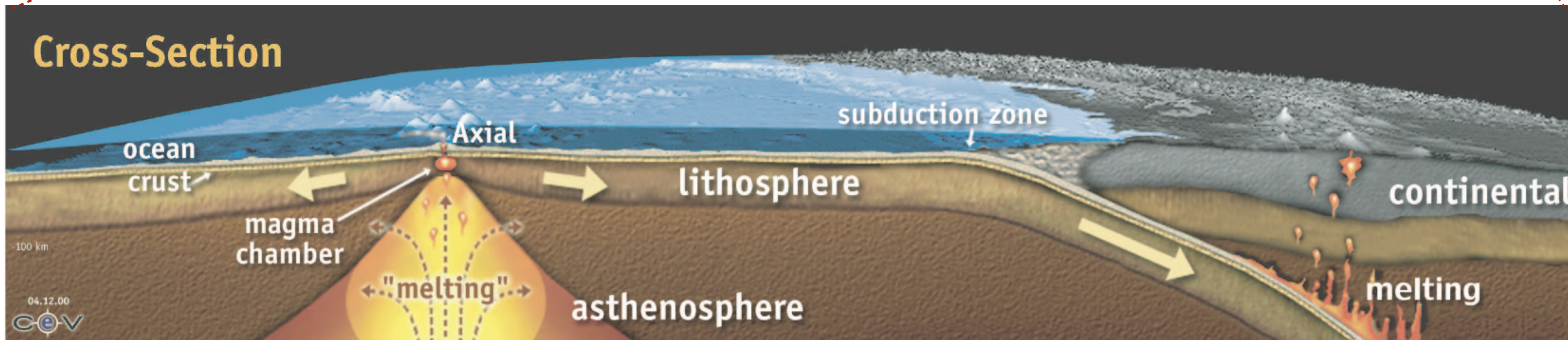
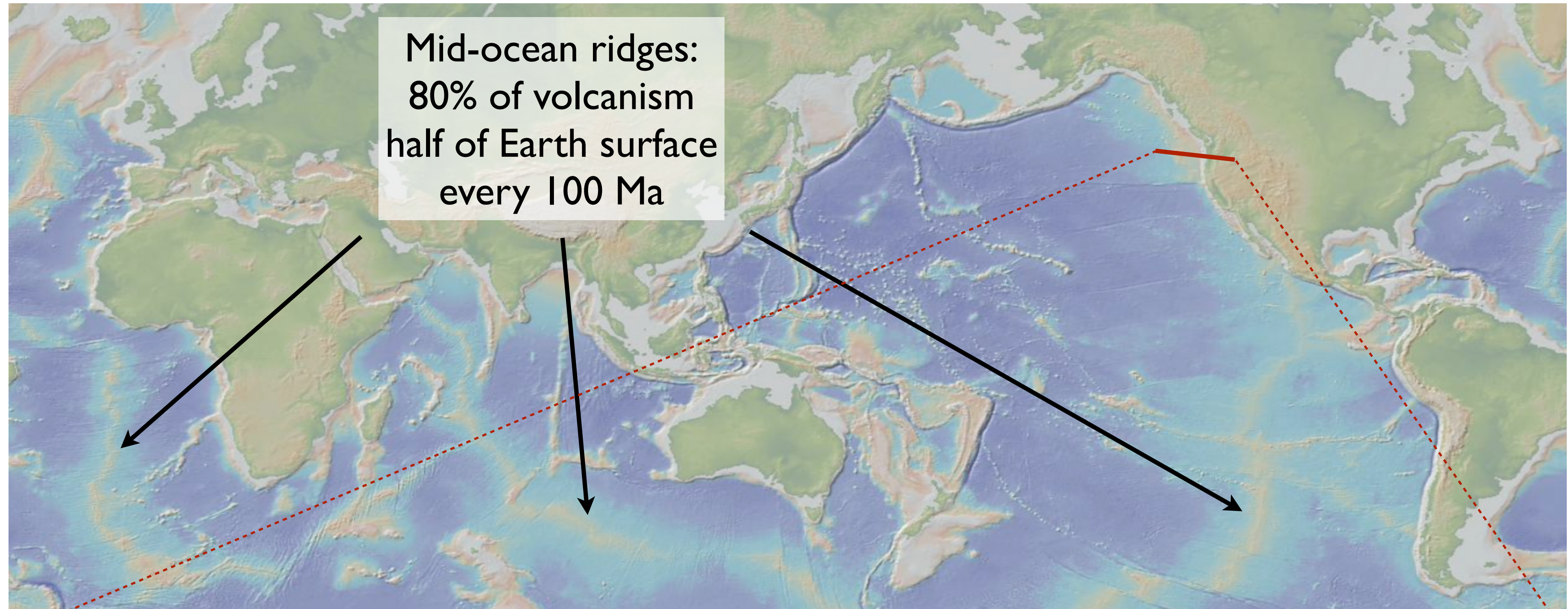
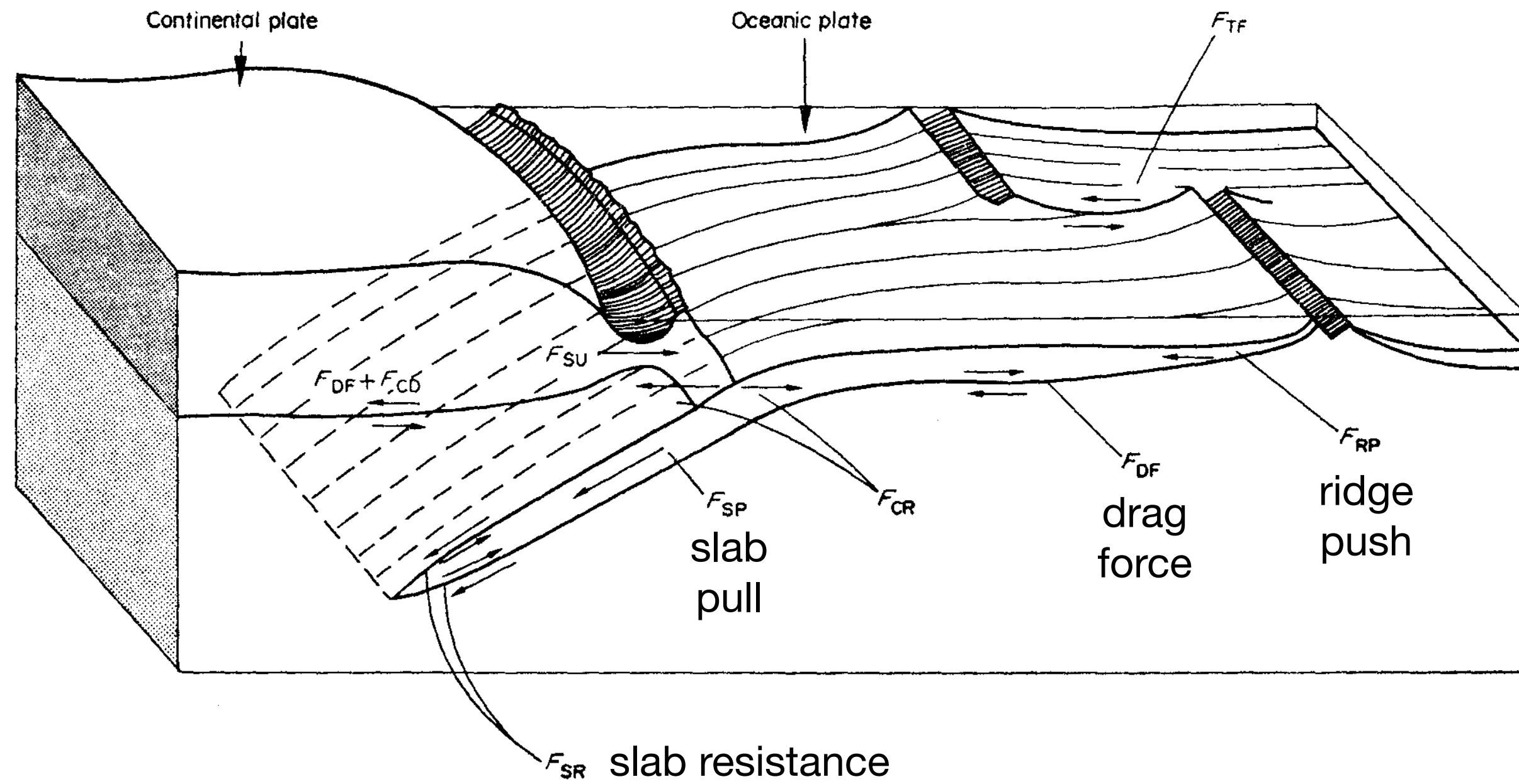


Plate tectonics controls sea-floor volcanism



Forces driving and resisting plate motion



Summary of the forces

Force	Form	Direction	Importance, X_s	Relative strength	Uncertainty	Units
F_{RP}	$\propto dl$	\perp strike	0.075	0.36	± 0.10	km^{-1}
F_{CR}	$\propto dl$	opp. rel. motion	0.040	0.16	± 0.09	km^{-1}
F_{TF}	$\propto dl$	opp. rel. motion	0.063	0.36	± 0.13	km^{-1}
F_{SP}	$\propto dl$	\perp strike	0.745	6.43	± 0.19	km^{-1}
F_{SU}	$\propto dl$	\perp strike	0.044	0.50	± 0.25	km^{-1}
F_{SR}	$\propto V_M \perp dl$	\perp strike	0.652	0.89	± 0.03	$\text{km}^{-1} \text{cm}^{-1} \text{yr}$
F_{CD}	$\propto V_M dA$	opp. abs. motion	0.056	5.65	± 2.22	$10^{-5} \text{km}^{-2} \text{cm}^{-1} \text{yr}$
F_{DF}	$\propto V_M dA$	opp. abs. motion	0.061	0.82	± 0.30	$10^{-5} \text{km}^{-2} \text{cm}^{-1} \text{yr}$

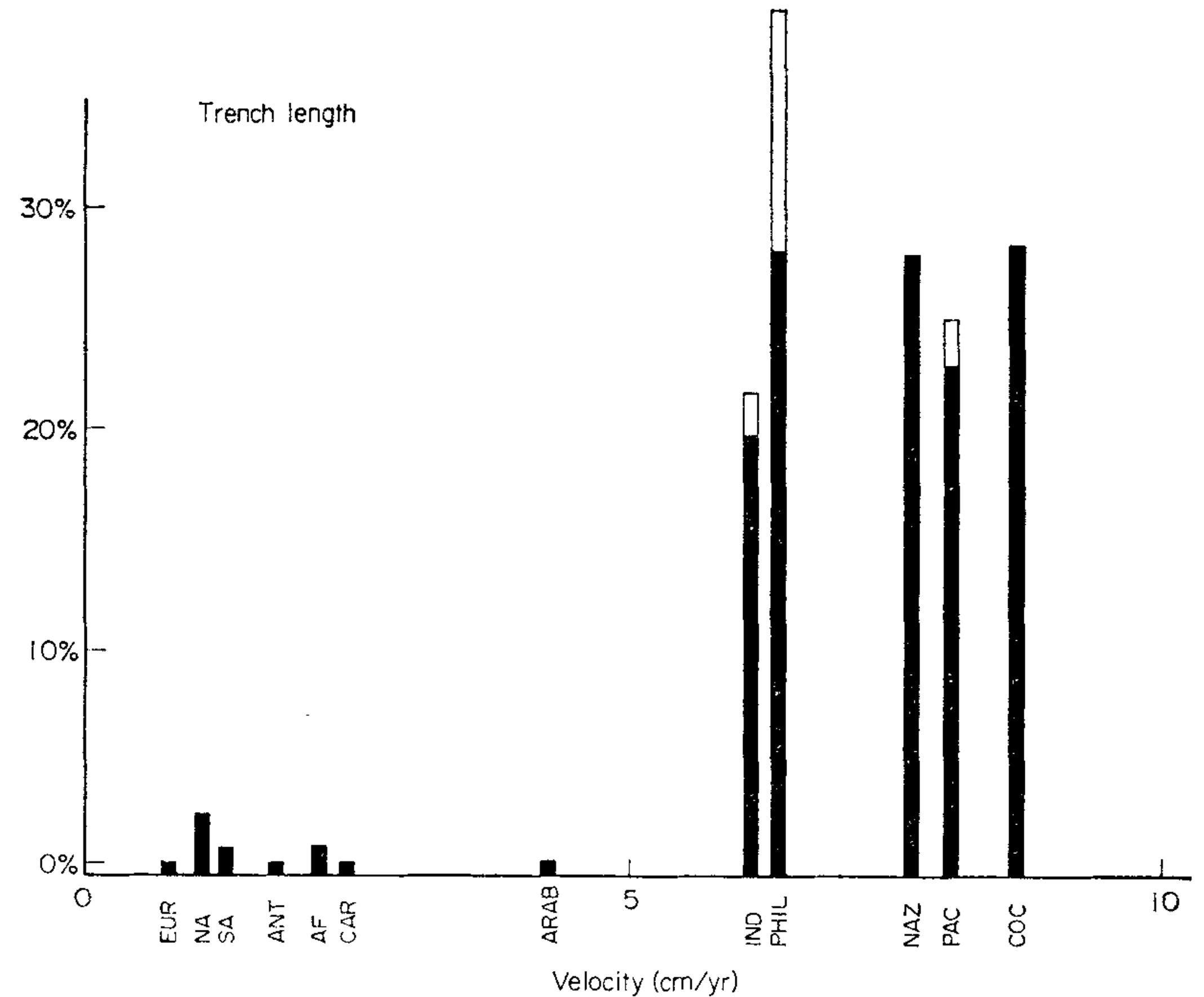
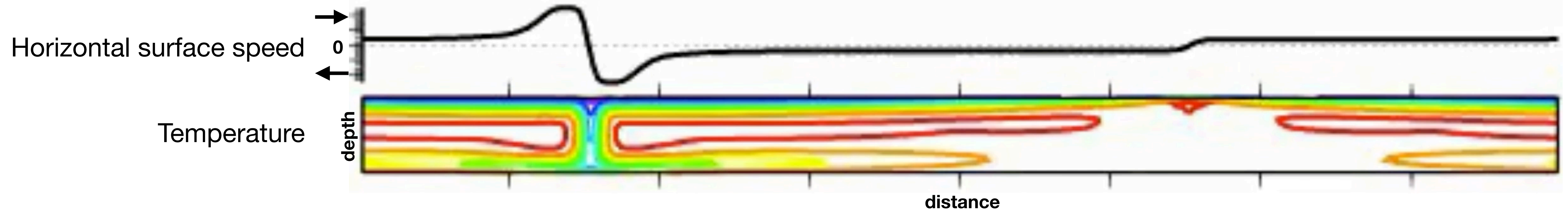


FIG. 8. Percentage of circumference of plate connected to downgoing slab. Open bar is total length, filled bar is effective length.

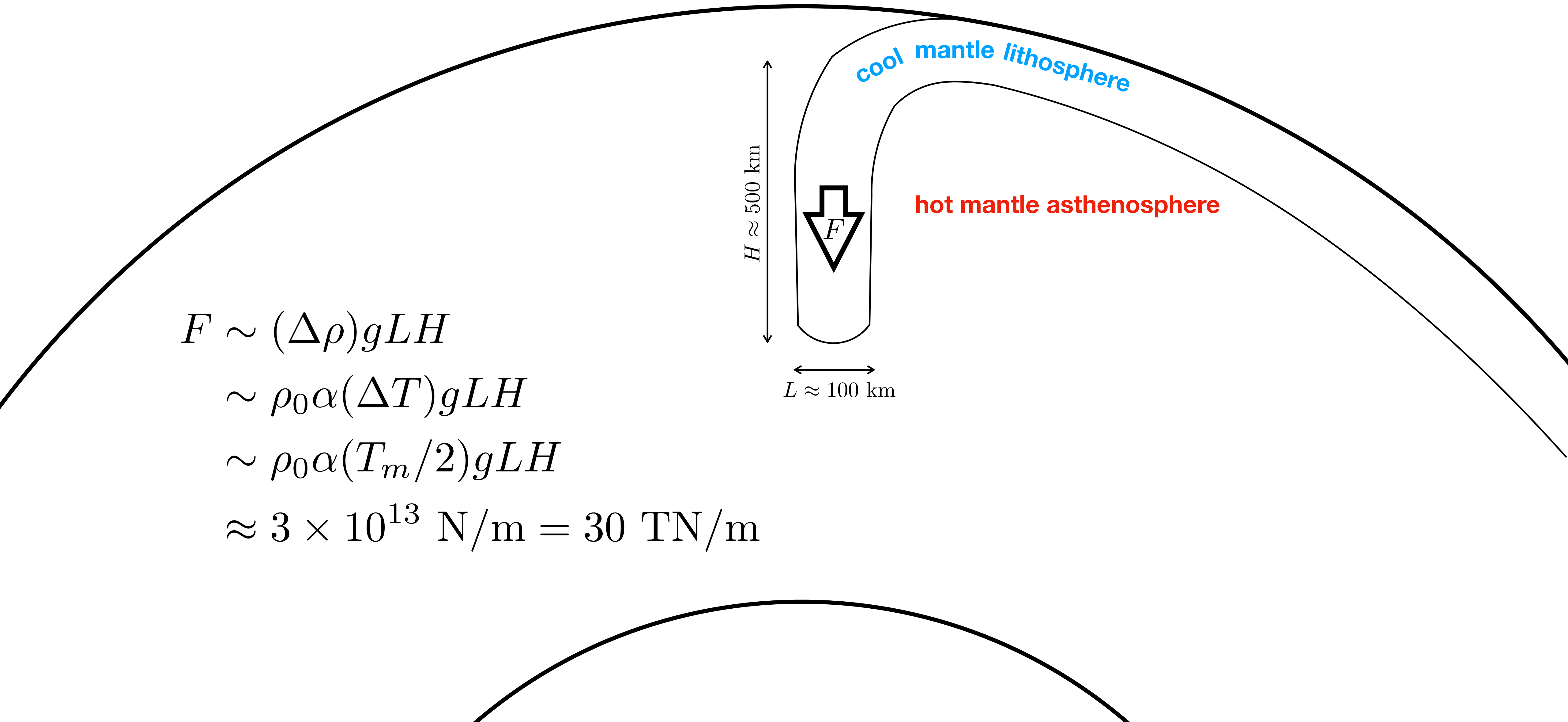
Internal heating + T-dependent viscosity + Plastic yield stress = Plate Tectonics



Grigne, Labrosse & Tackley 2005

- broad plates with narrow boundaries
 - passive divergent boundaries (MORs)
 - active convergent boundaries (SZs)
- physical basis(es) for yield stress
 - chemical heterogeneity
 - continents and continental rifting
 - one-sided subduction
 - strike-slip (poloidal) motion
 - role of magma in mechanics

Force of slab pull – simple estimate



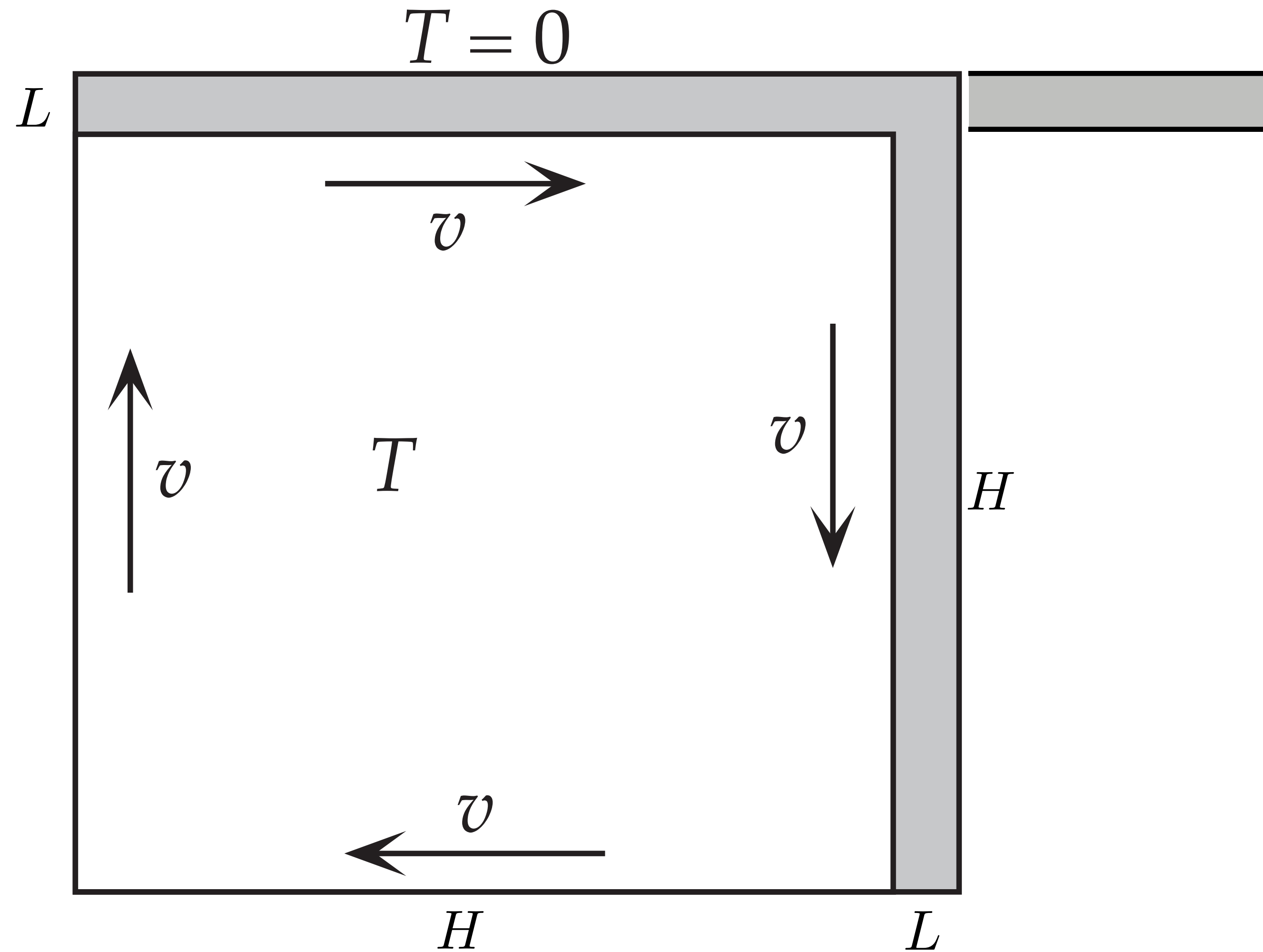
$$F \sim (\Delta\rho)gLH$$

$$\sim \rho_0\alpha(\Delta T)gLH$$

$$\sim \rho_0\alpha(T_m/2)gLH$$

$$\approx 3 \times 10^{13} \text{ N/m} = 30 \text{ TN/m}$$

Slab pull and plate speed — simple estimate

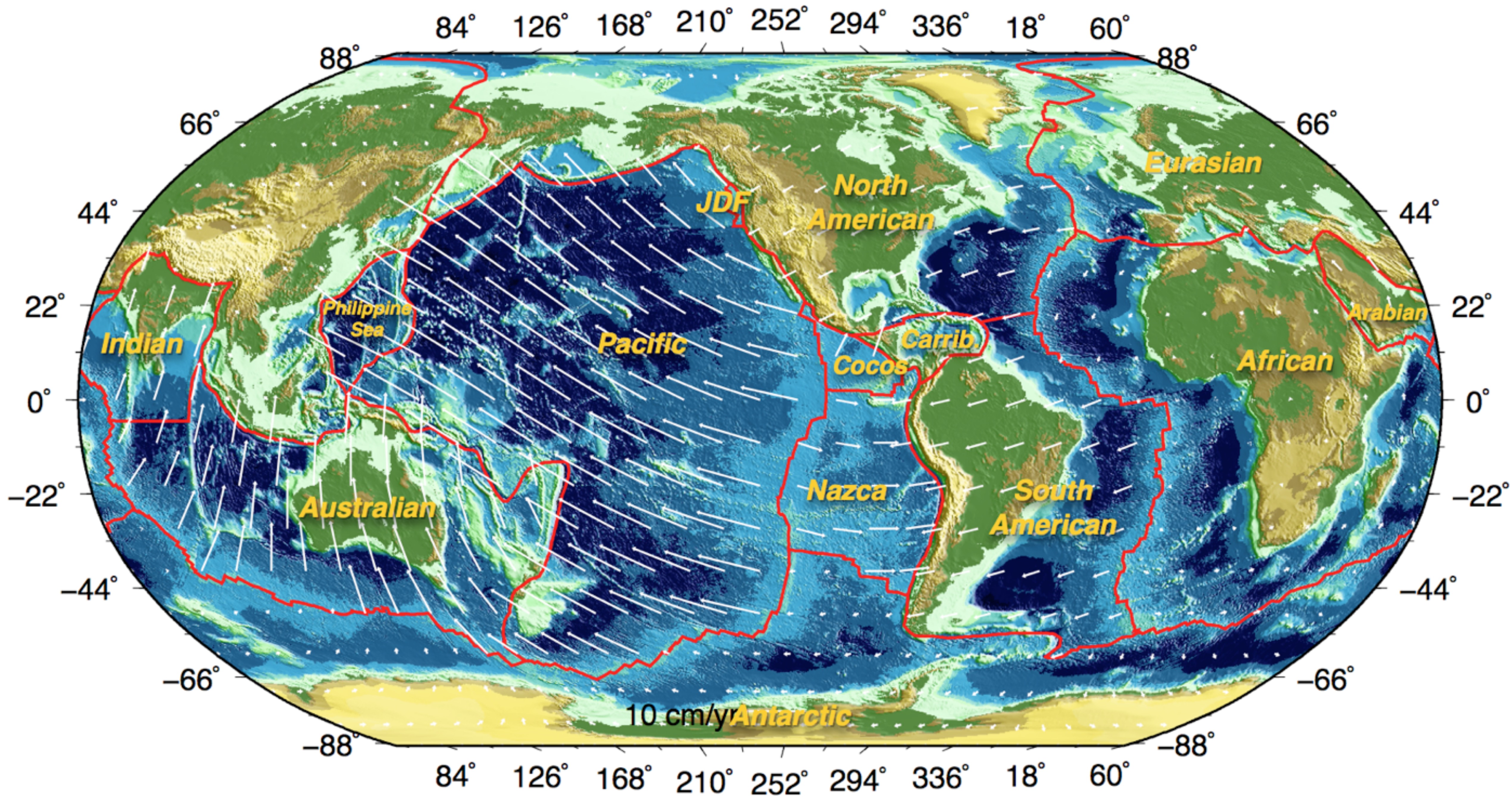


Newton's 2nd law: $F_D + F_R \approx 0$

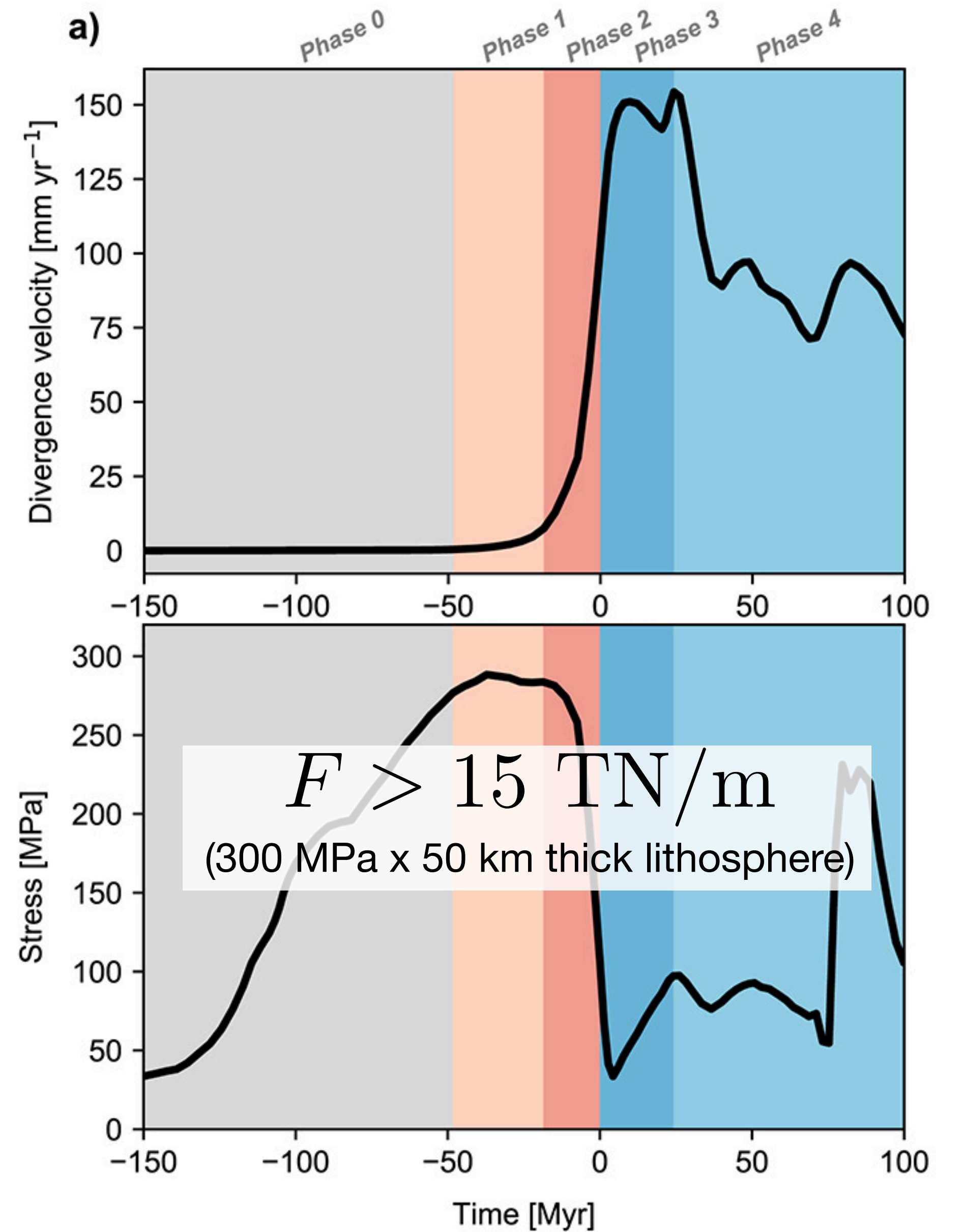
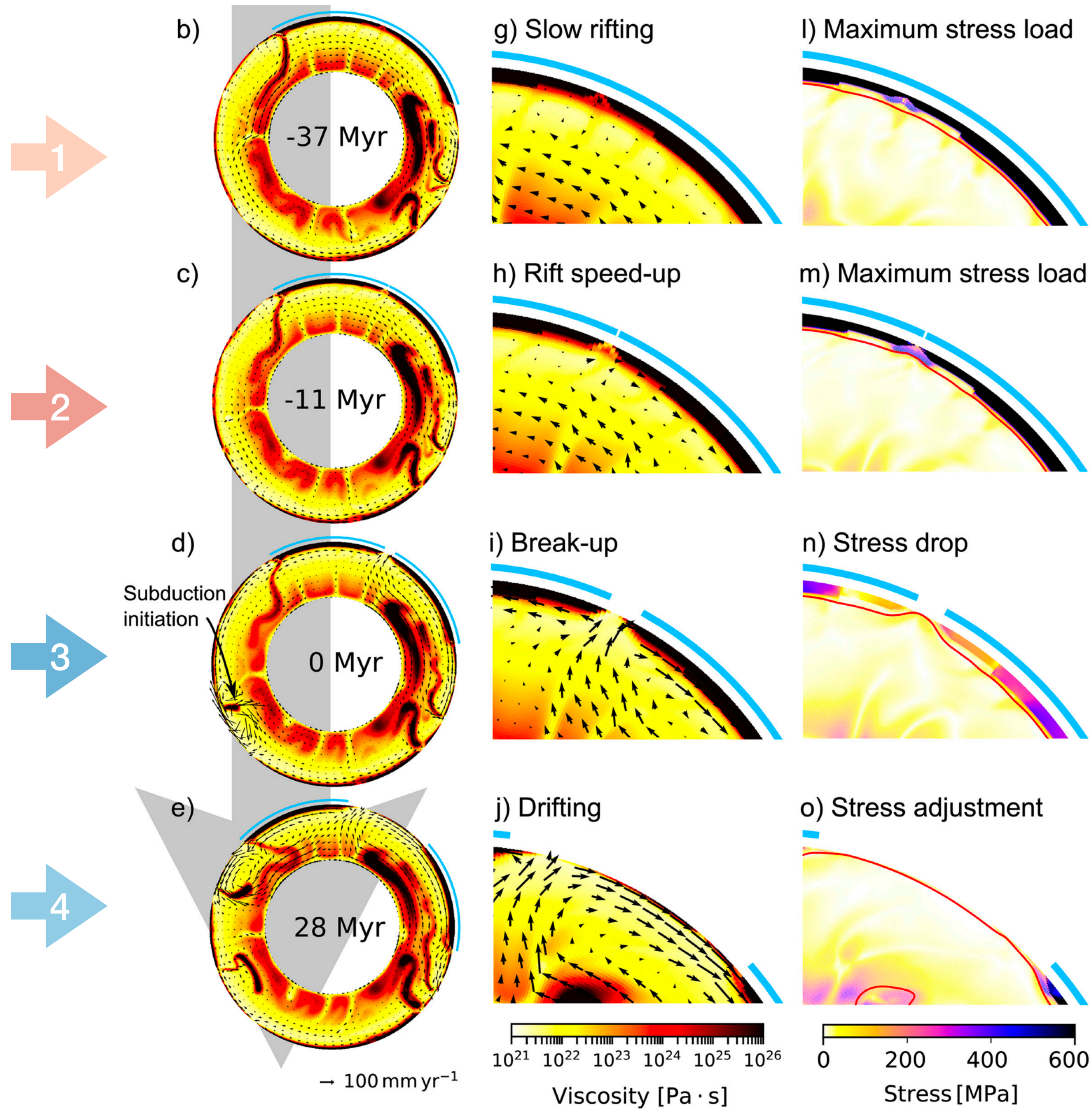
$$F_D \sim \rho_0 \alpha (T_m / 2) g L H$$

$$F_R \sim 3H (\eta \times 2v / H)$$

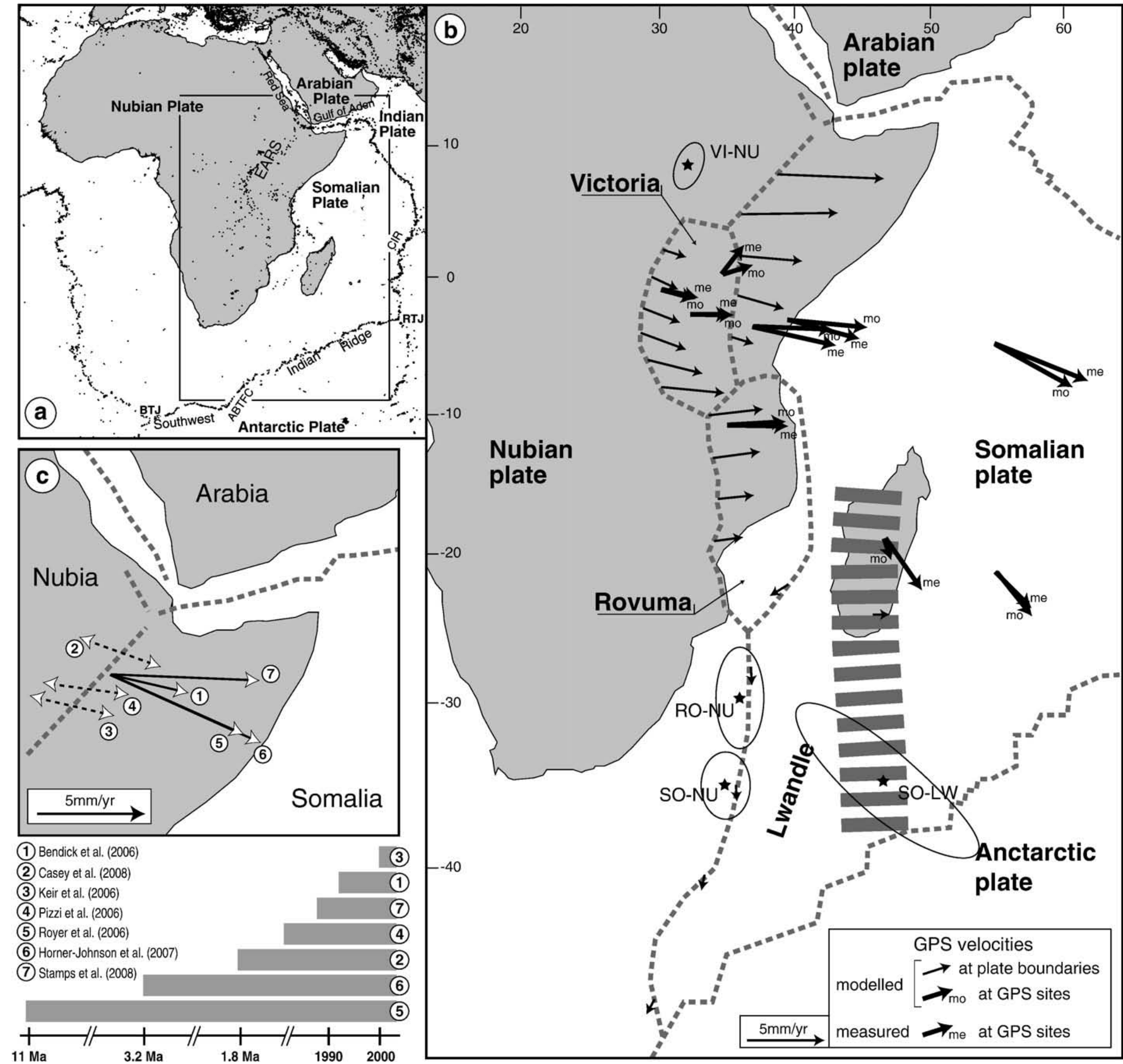
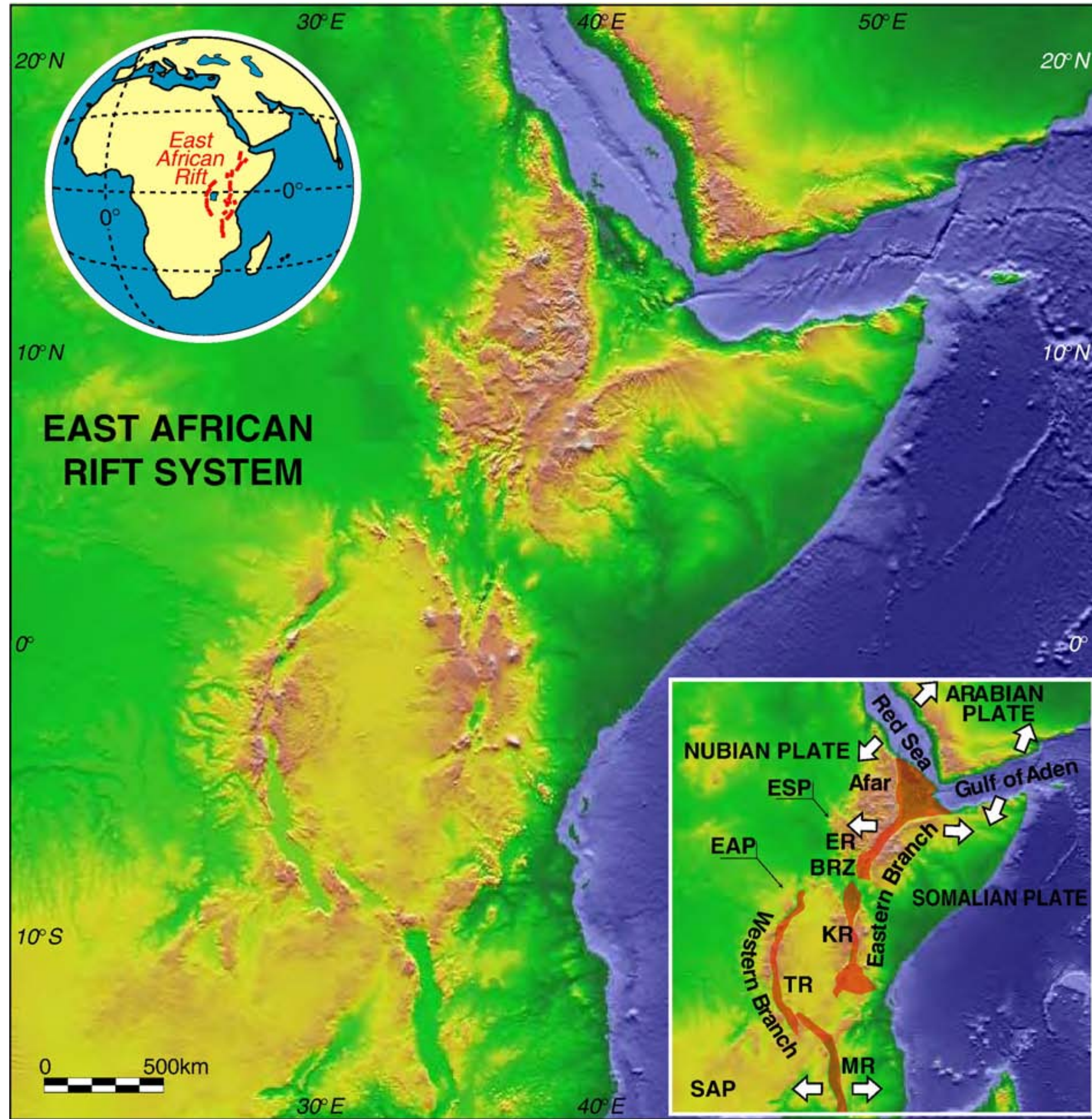
$$\Rightarrow v = \rho_0 \alpha T g L H / 12 \eta$$
$$\approx 10 \text{ cm/yr}$$



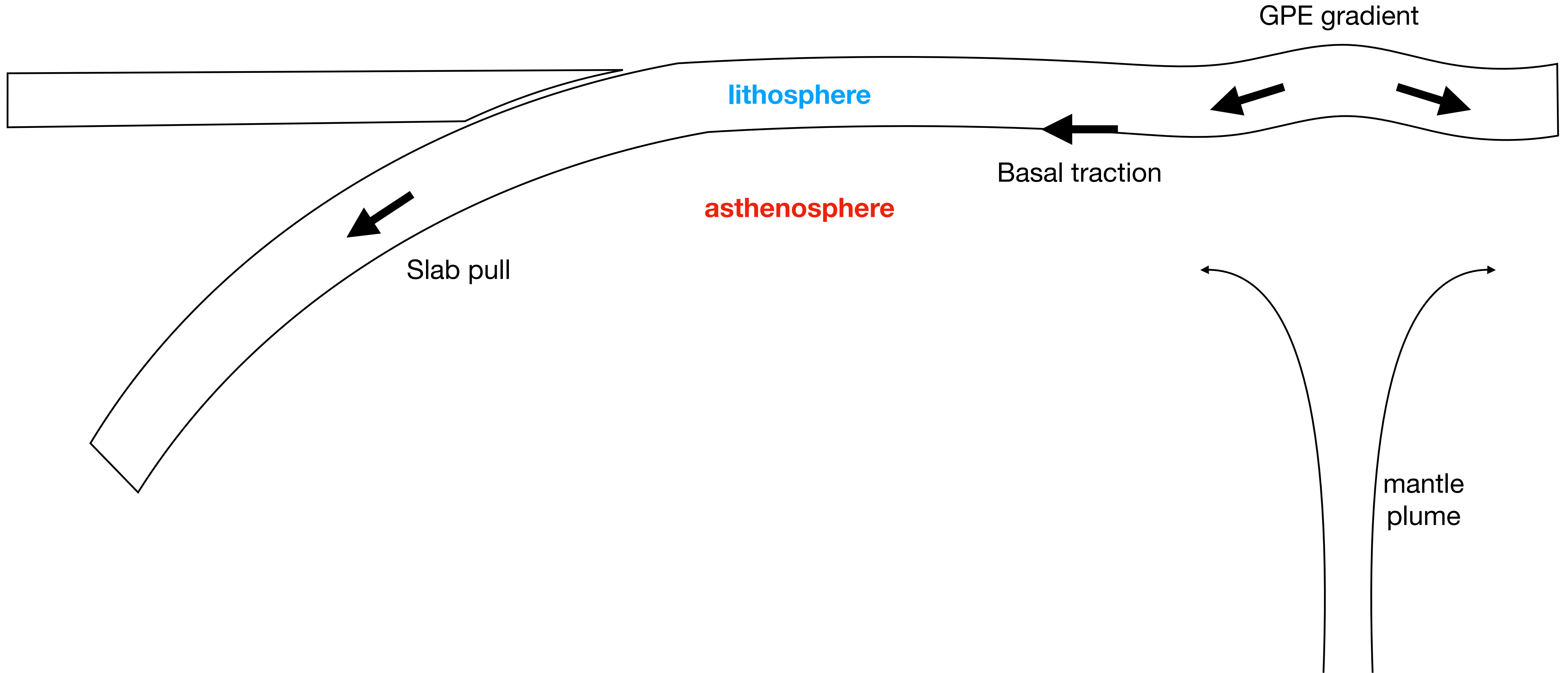
Force within lithosphere causing rifting



What is happening here in Africa?



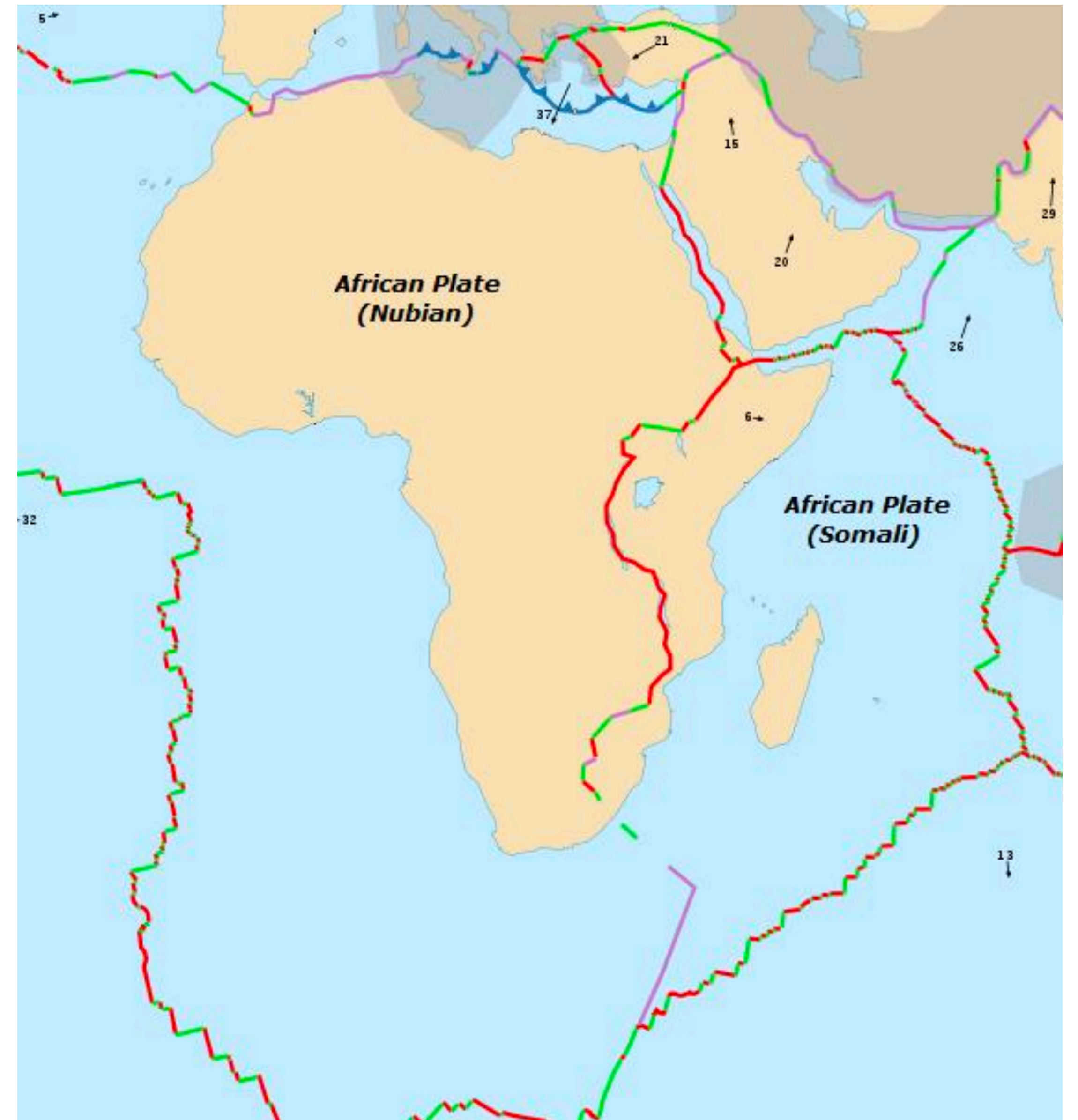
Candidate forces for rifting continents



What is happening here in Africa?

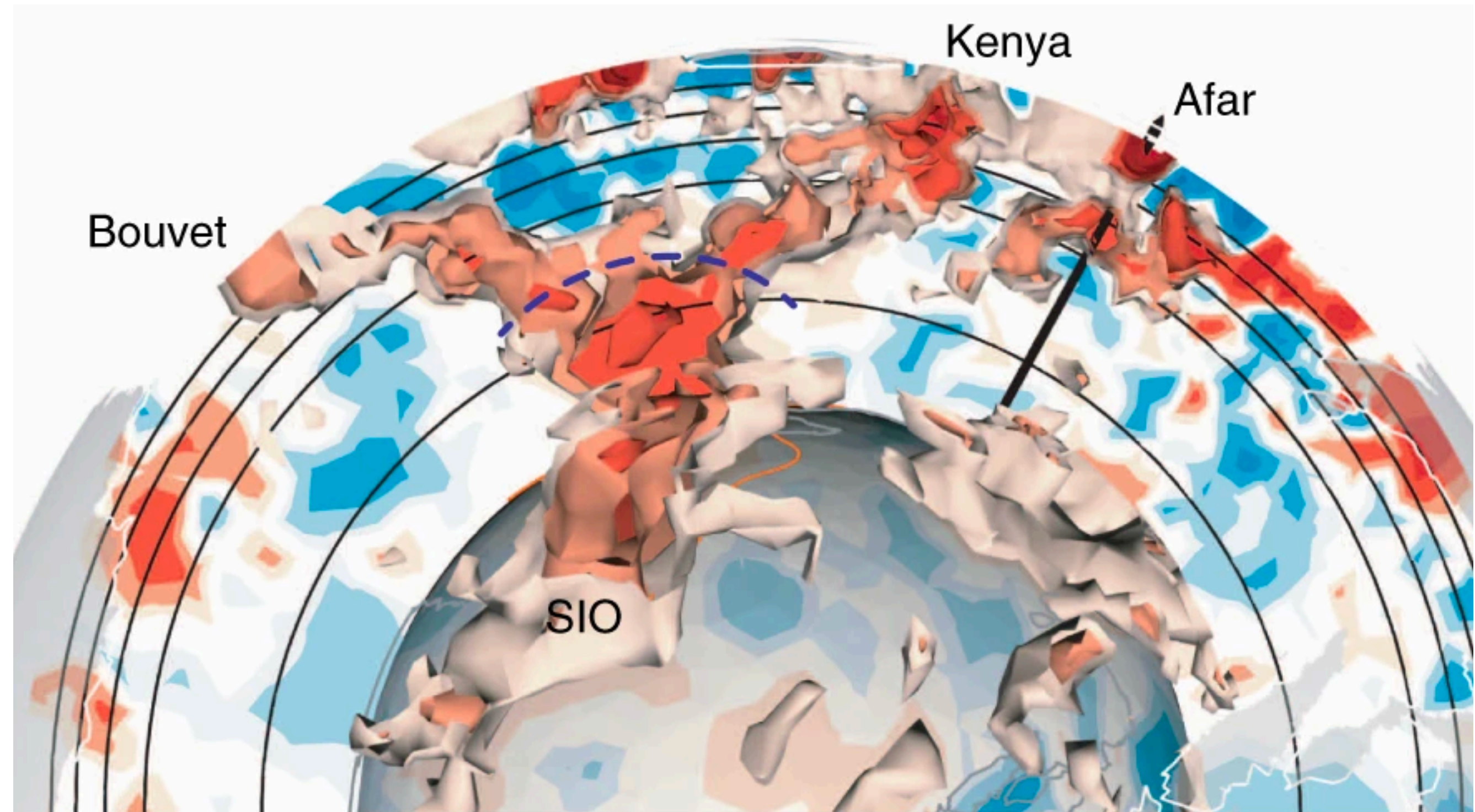
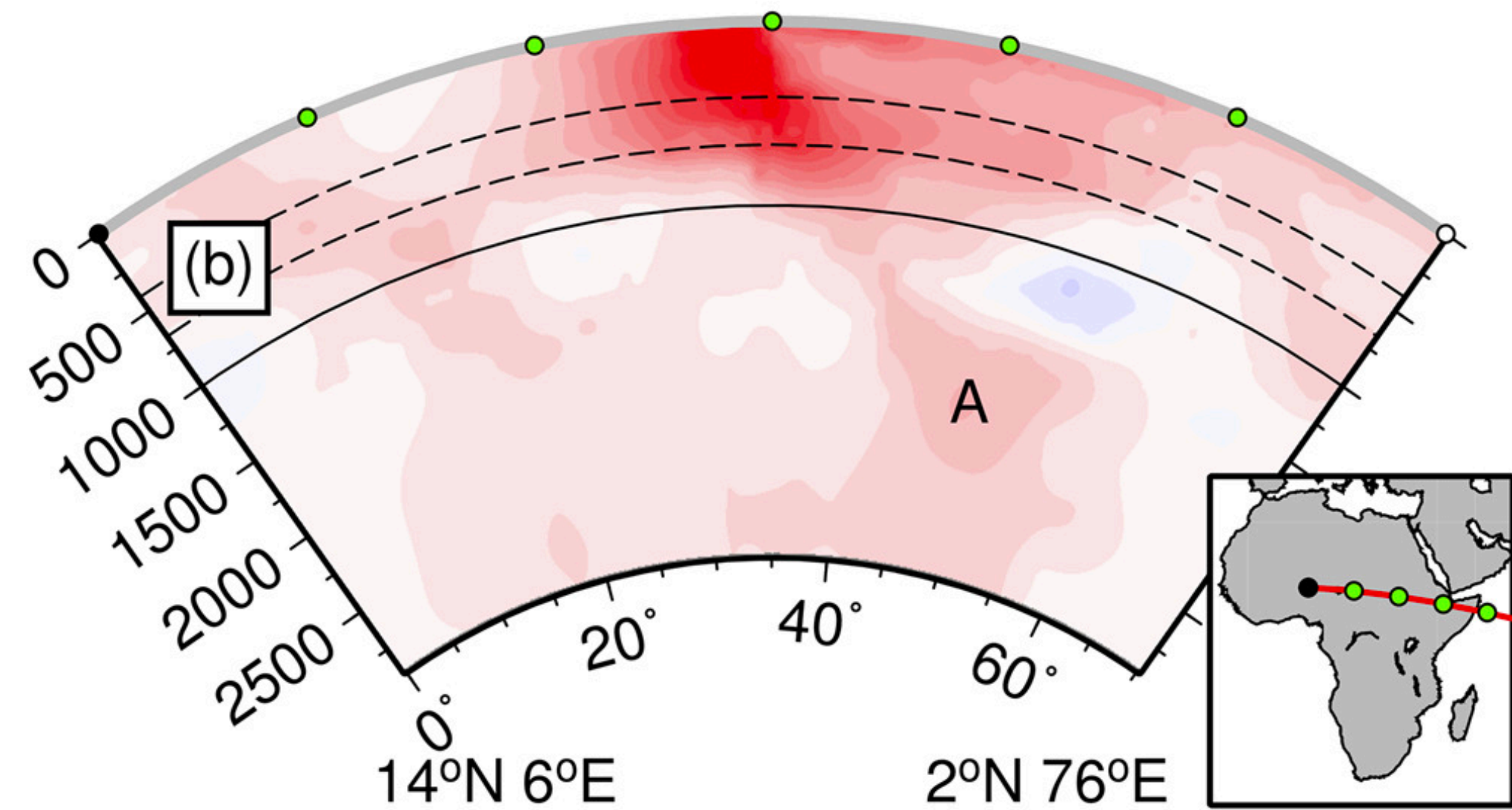
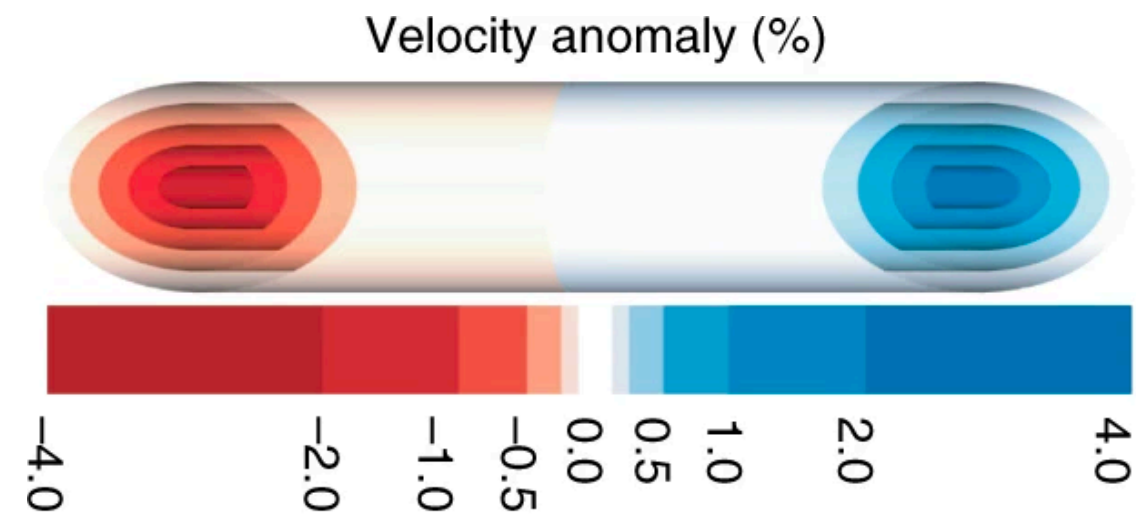
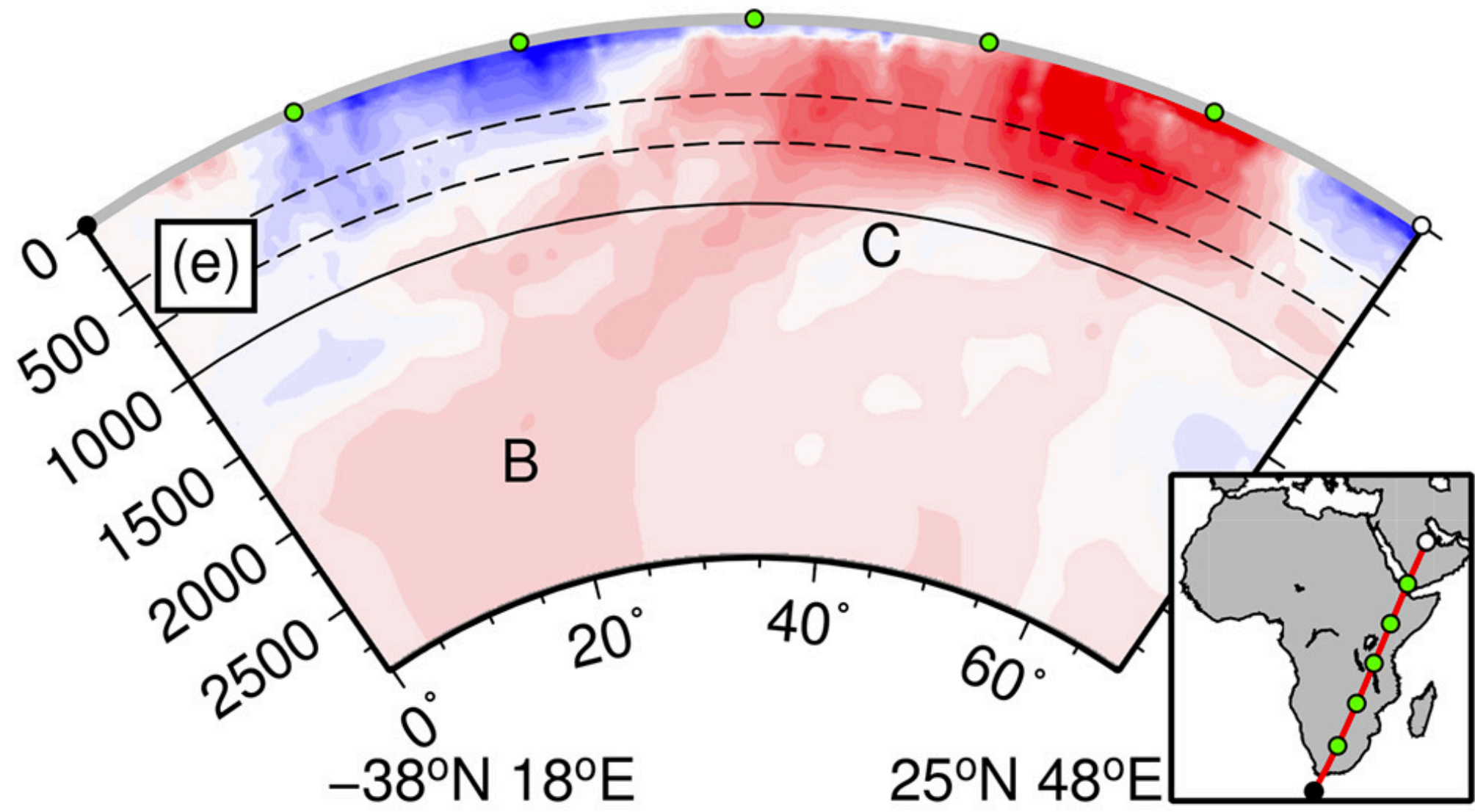


Heezen & Tharp, *PTRSA* 1965

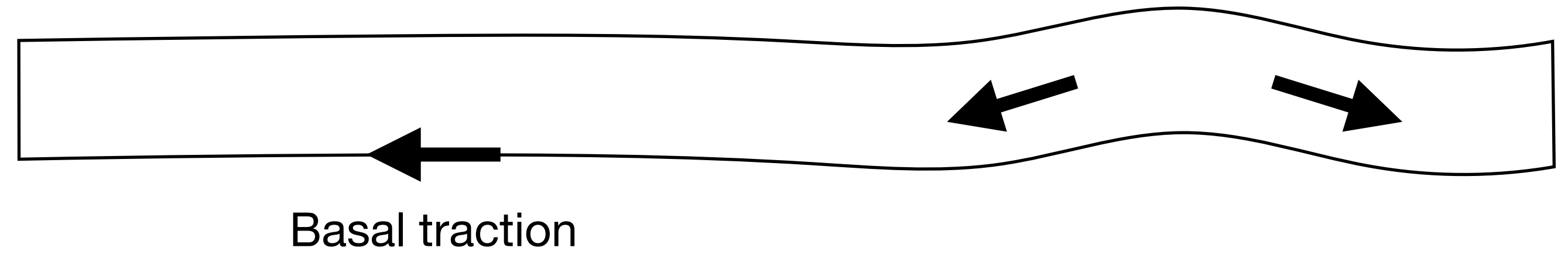


McGill University

What is happening here in Africa?



Basal traction force (from mantle convection)

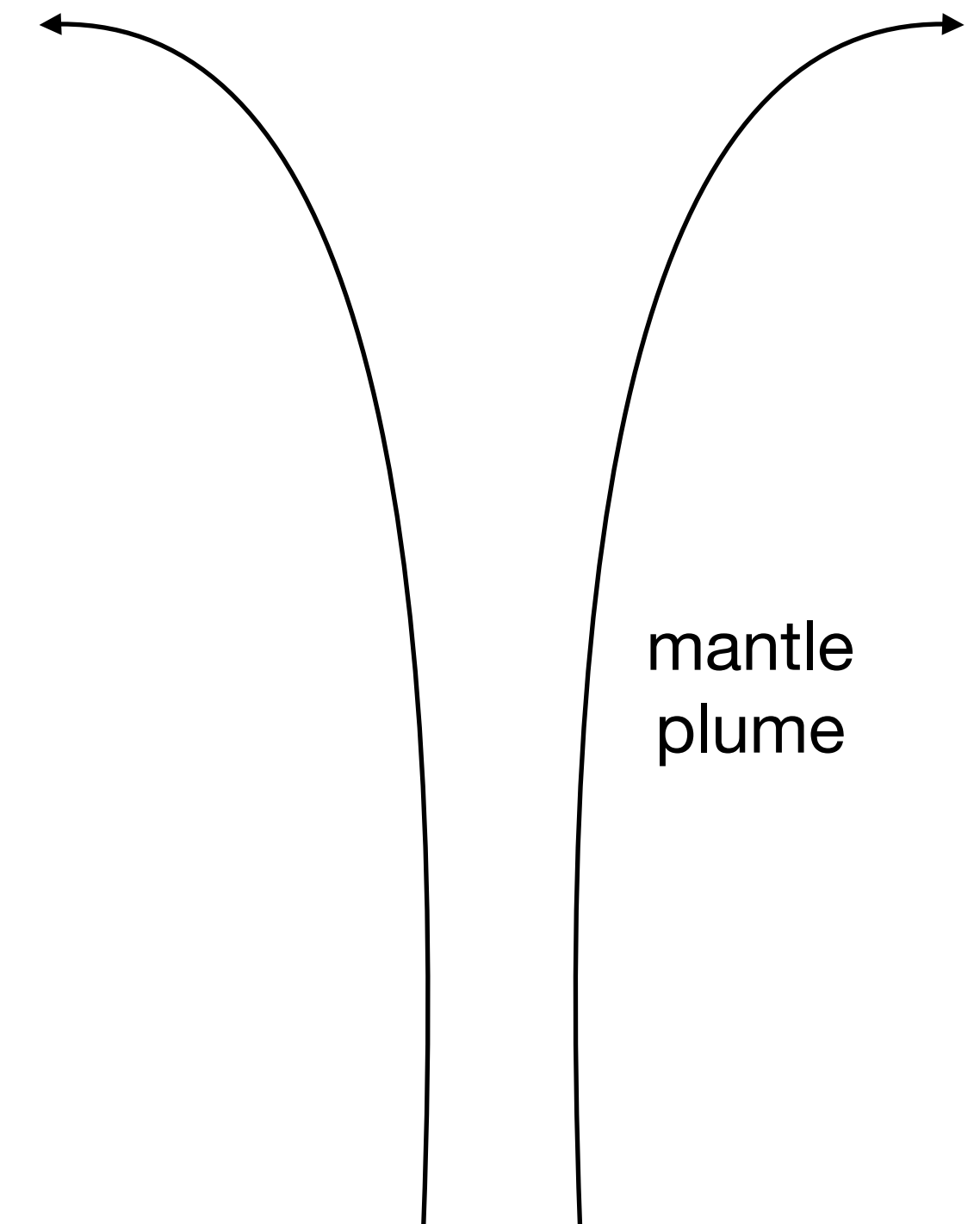
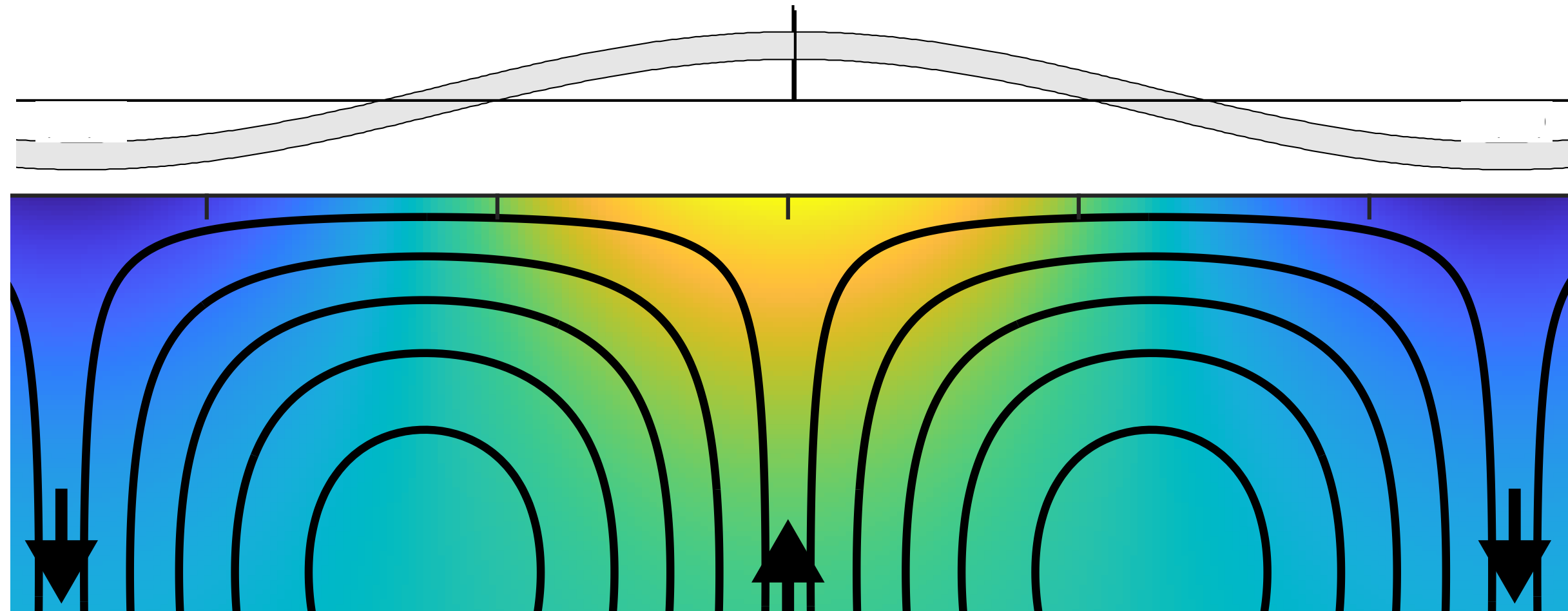
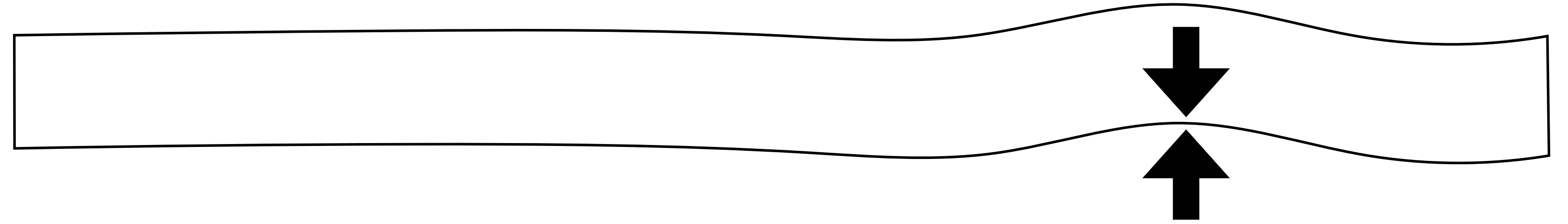


$$F \leq 3 \text{ MPa} \times 1000 \text{ km} = 3 \times 10^{12} \text{ N/m}$$

typical convective stress plate application length

Dynamic topography

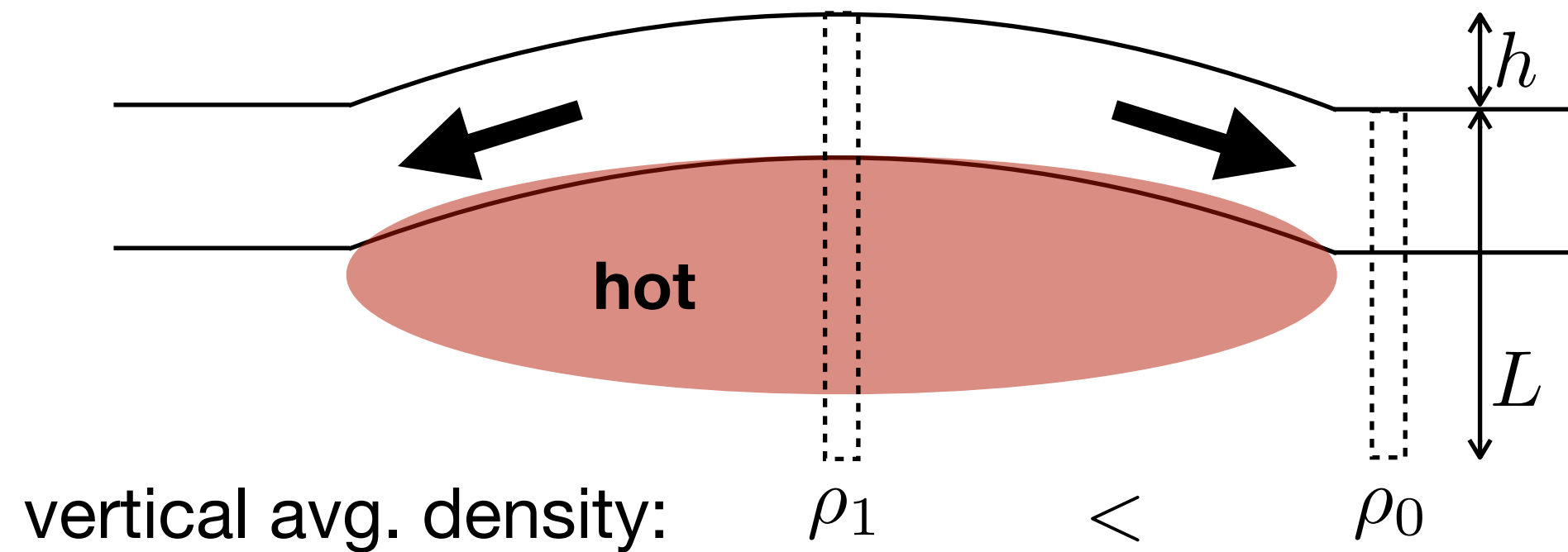
$$F_{\text{buoyancy}} + F_{\text{viscous}} - F_{\text{weight}} = 0$$



$$T(x, z) = T_m + \Delta T \cos(kx) \exp(kz)$$

$$\Delta h = \frac{3\rho_0\alpha\Delta T}{4\Delta\rho k} \cos(kx)$$

GPE gradient force (~same as ridge push force)



Pratt isostasy: $\rho_1(L + h) = \rho_0 L$

Potential energy:
$$\text{GPE}(x) = \int_{-L}^{h(x)} \rho(z)gz \, dz$$

Force per length:
$$F = \Delta\text{GPE} = \frac{\rho_1 g(L + h)^2}{2} - \frac{\rho_0 g L^2}{2}$$
$$= \frac{\rho_0 g L h}{2}$$

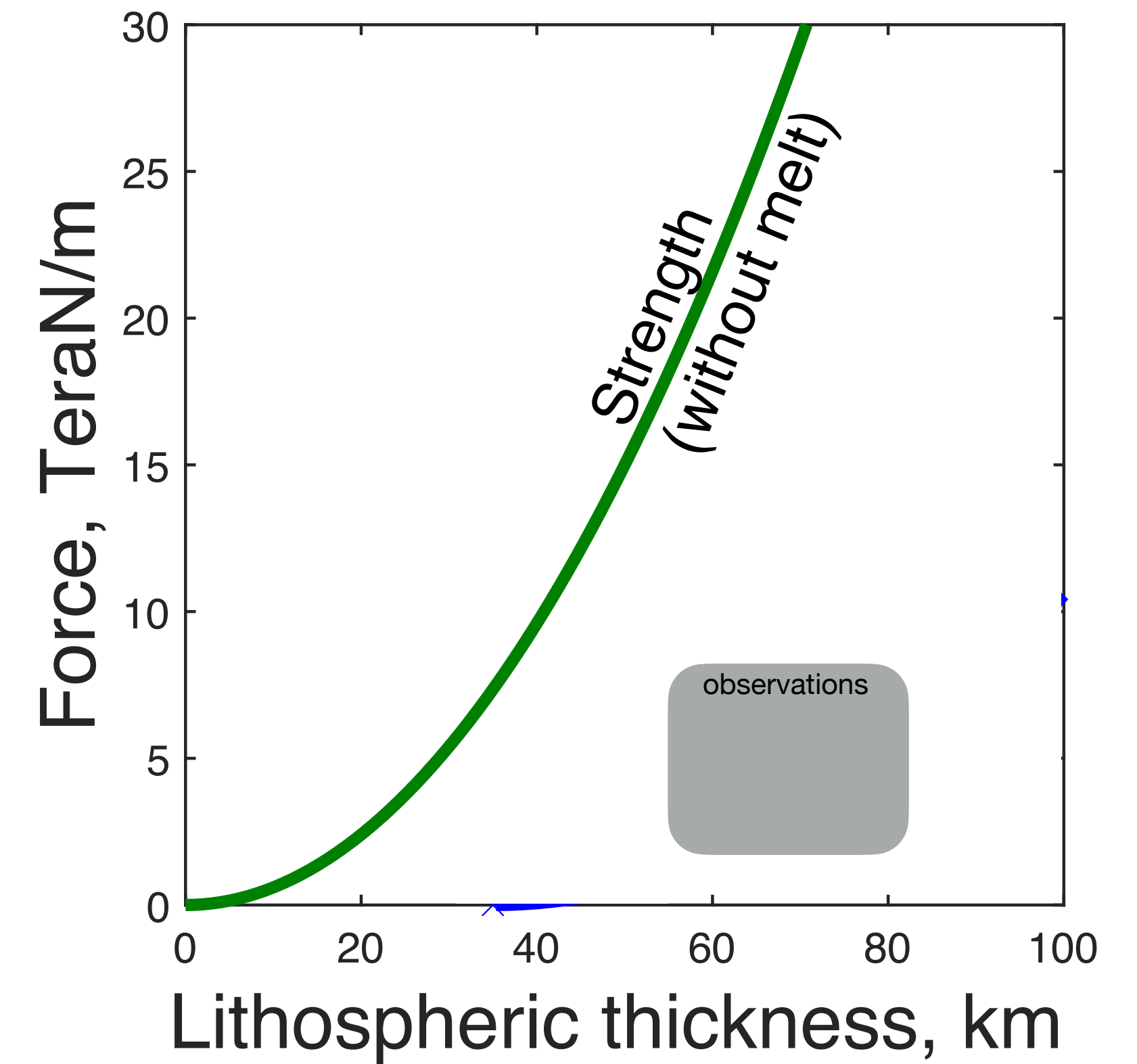
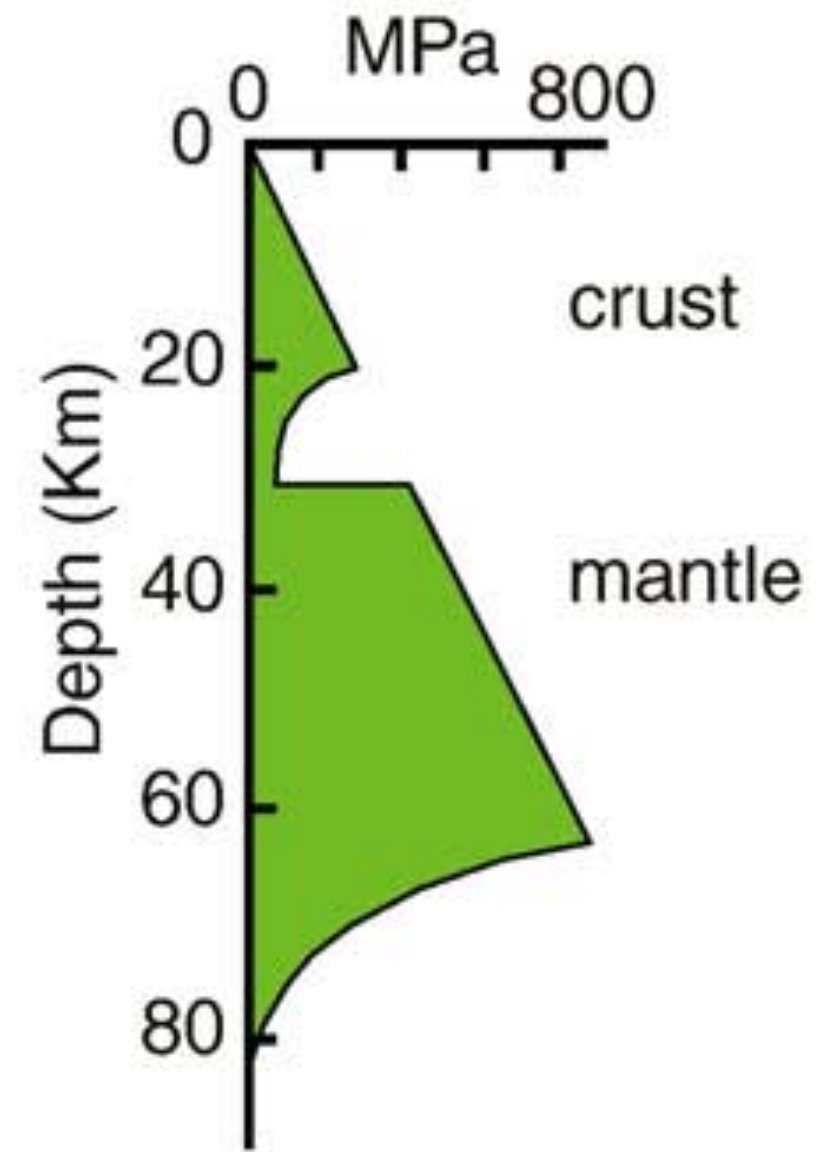
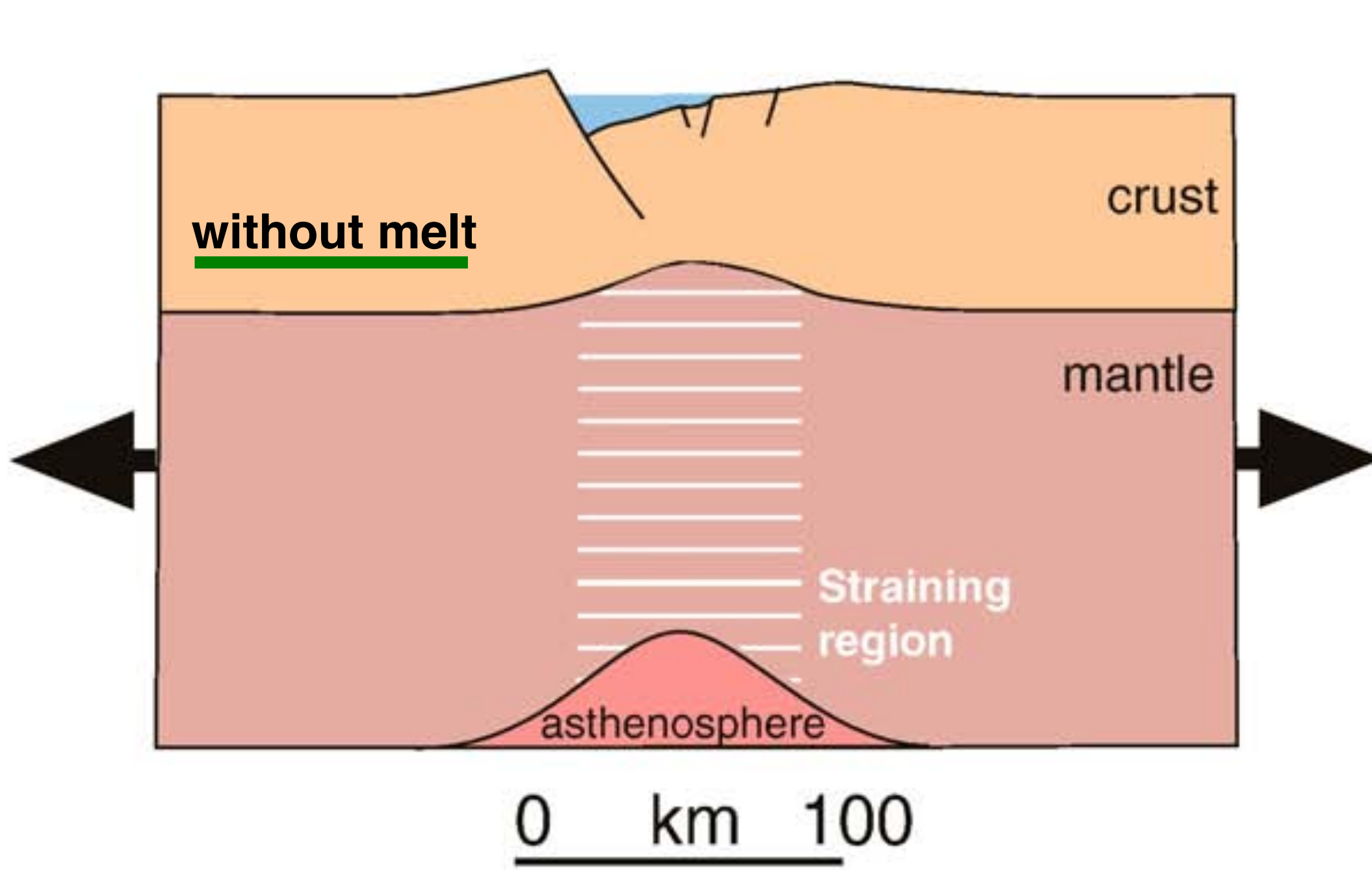
For an upper bound:

$$h = 1 \text{ km}, L = 150 \text{ km}, \rho = 3000 \text{ kg/m}^3$$

$$F \leq 5 \times 10^{12} \text{ N/m}$$

Observation: rifting can occur at driving force less than $\sim 5 \text{ TN/m}$

Rifting by frictional slip on faults



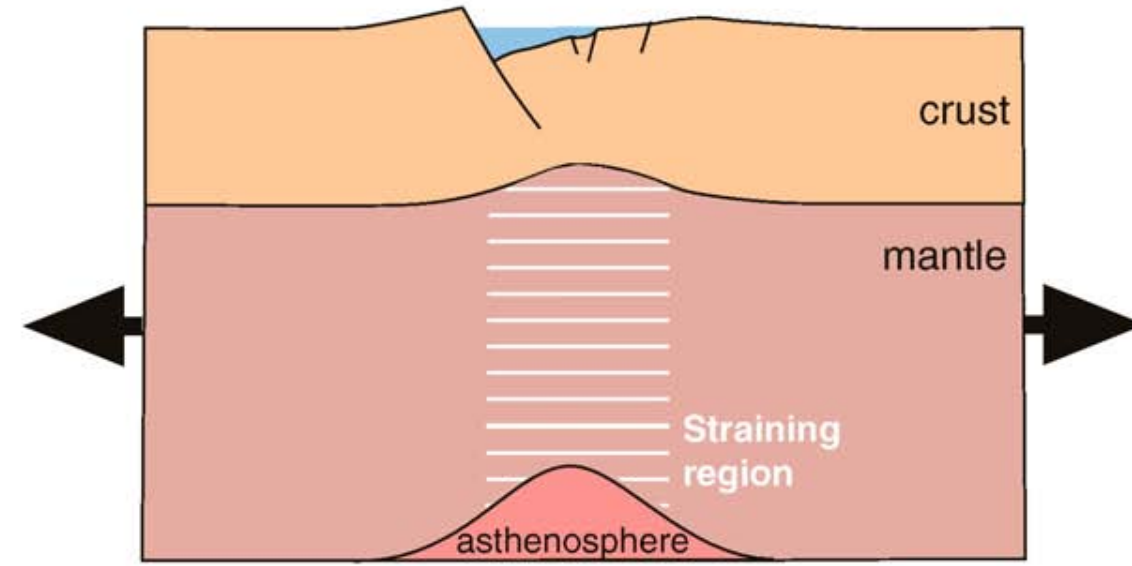
$$F \sim A(\rho_{\text{lith}} - \rho_w)gL^2$$

Andersonian
faulting strength

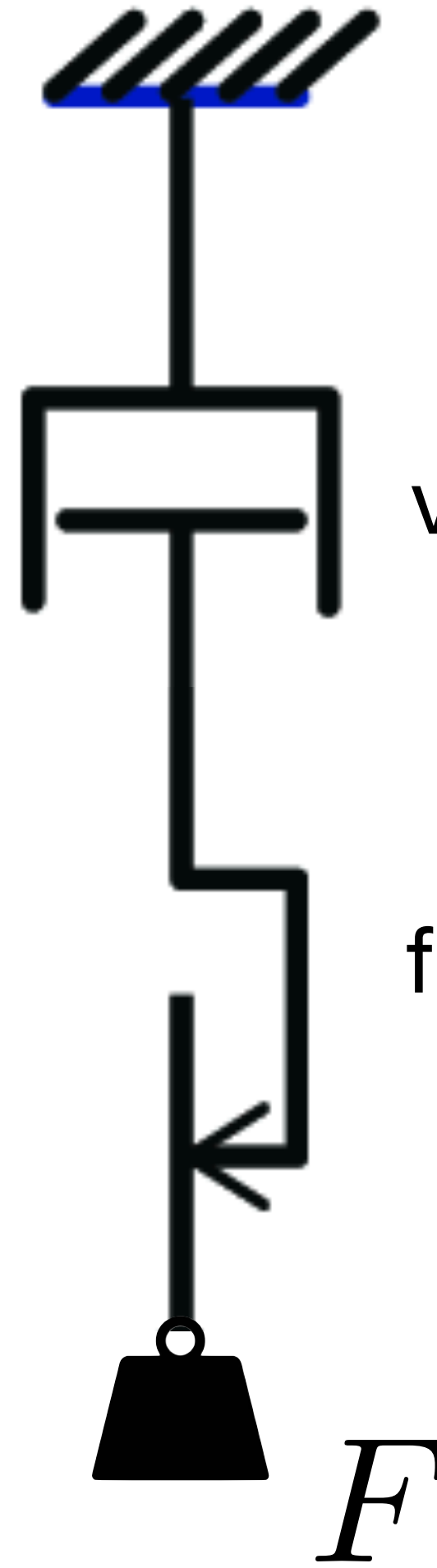
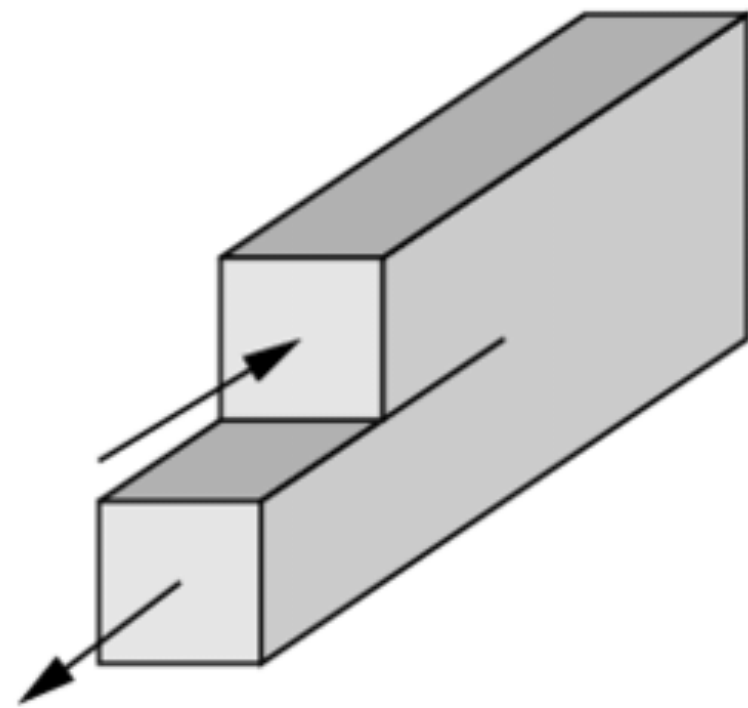
Hydrostatic
pressure

(N.B., $F = \sigma L$)

Using plasticity to represent faults



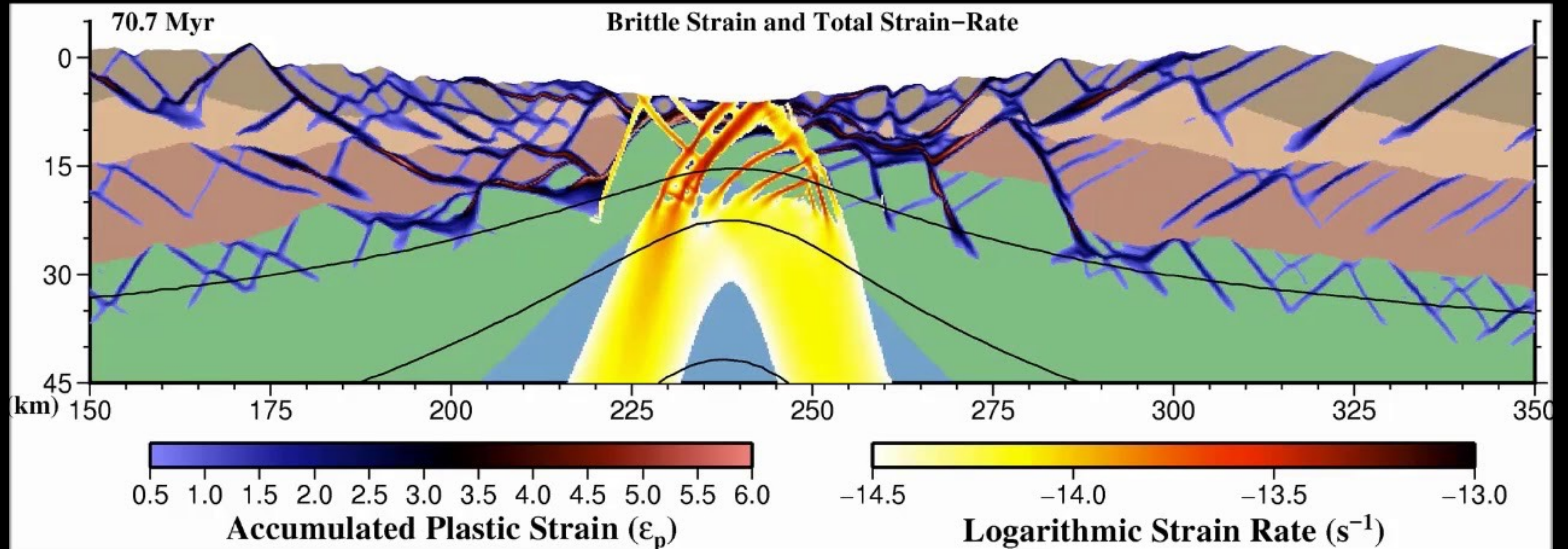
Mode II



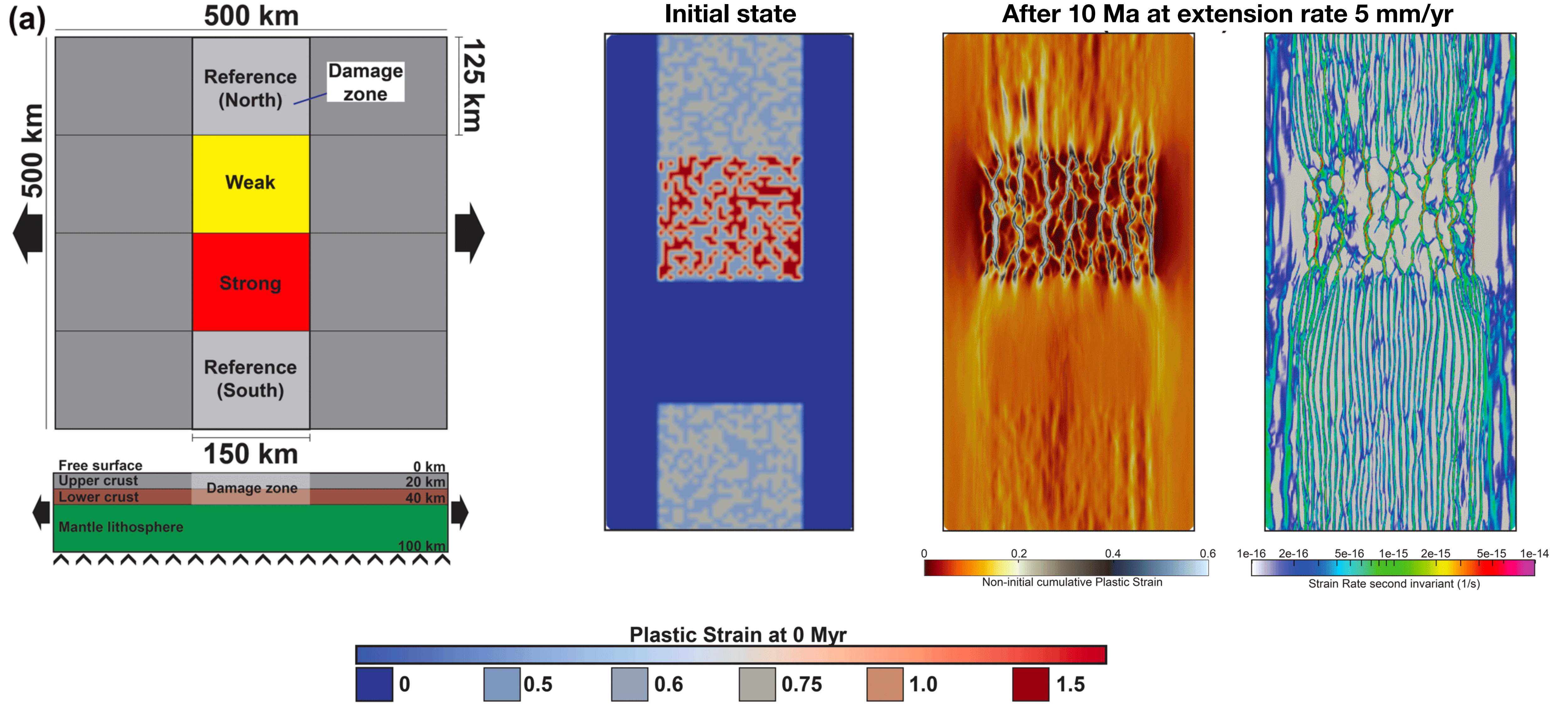
viscous dashpot representing *creep deformation*
extends at a rate $\propto F$

frictional slider representing *fracture and fault slip*
extends if $F > F_{\text{crit}}$

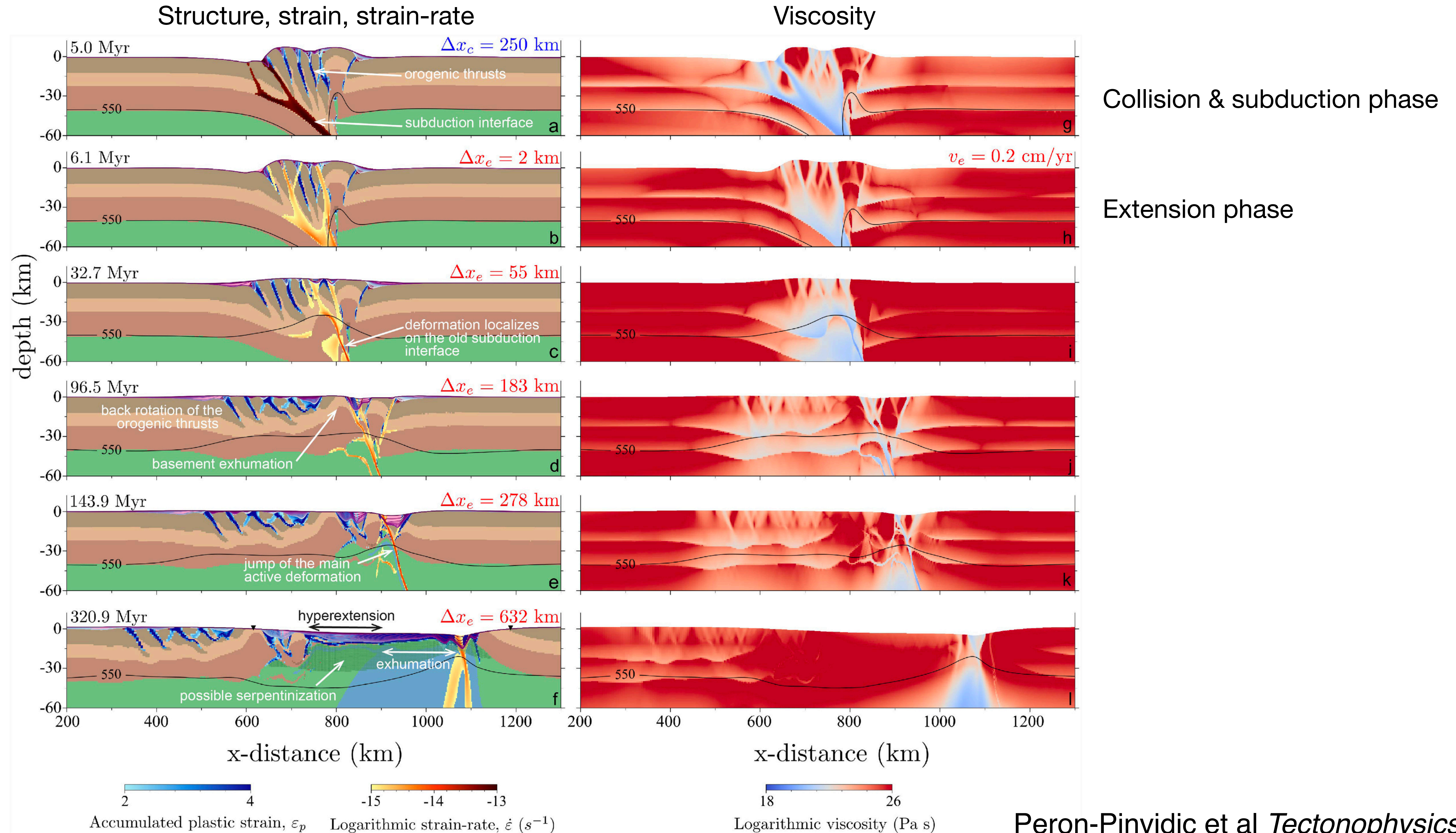
Faulting & frictional slip modelled as plastic deformation



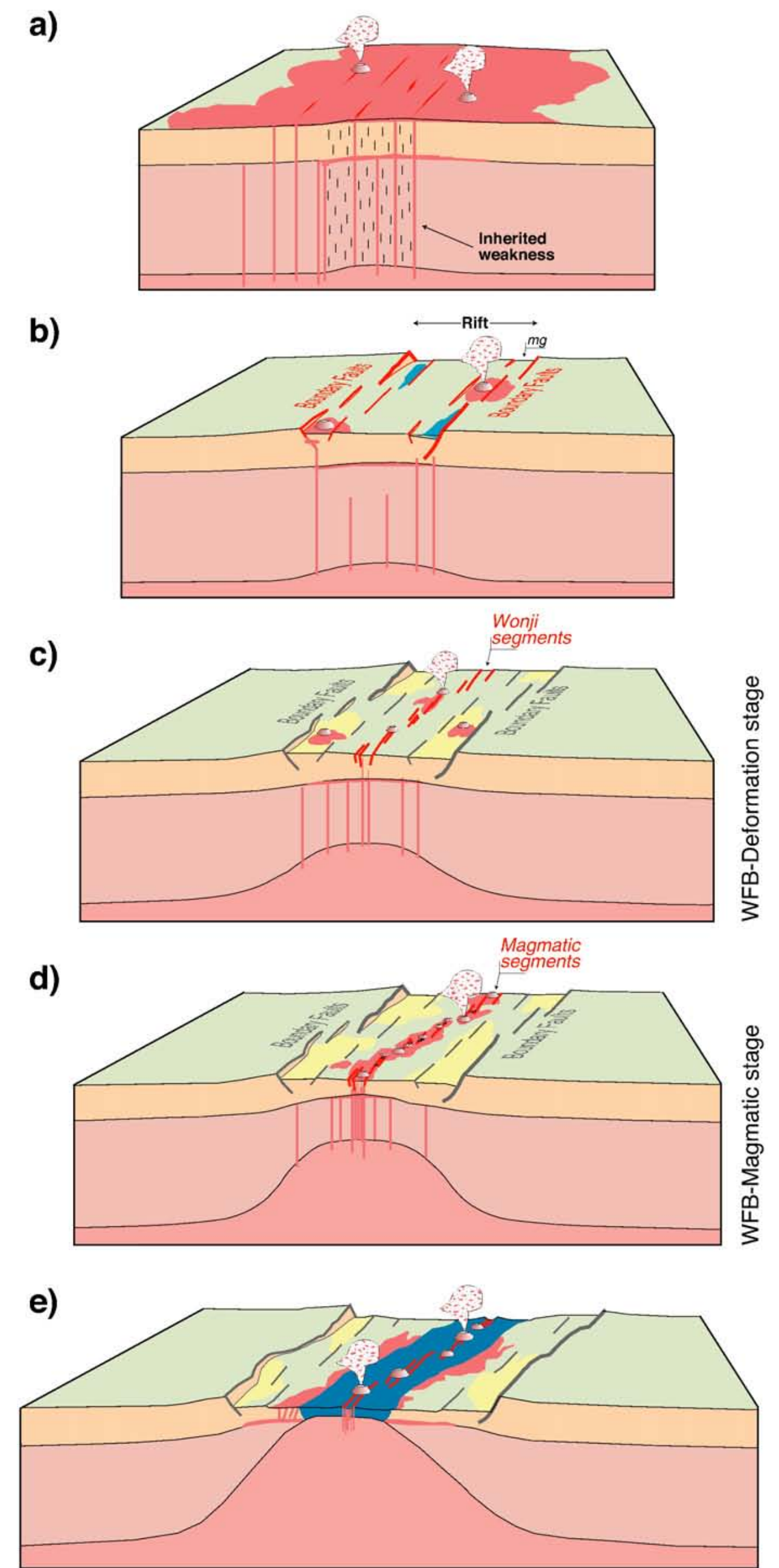
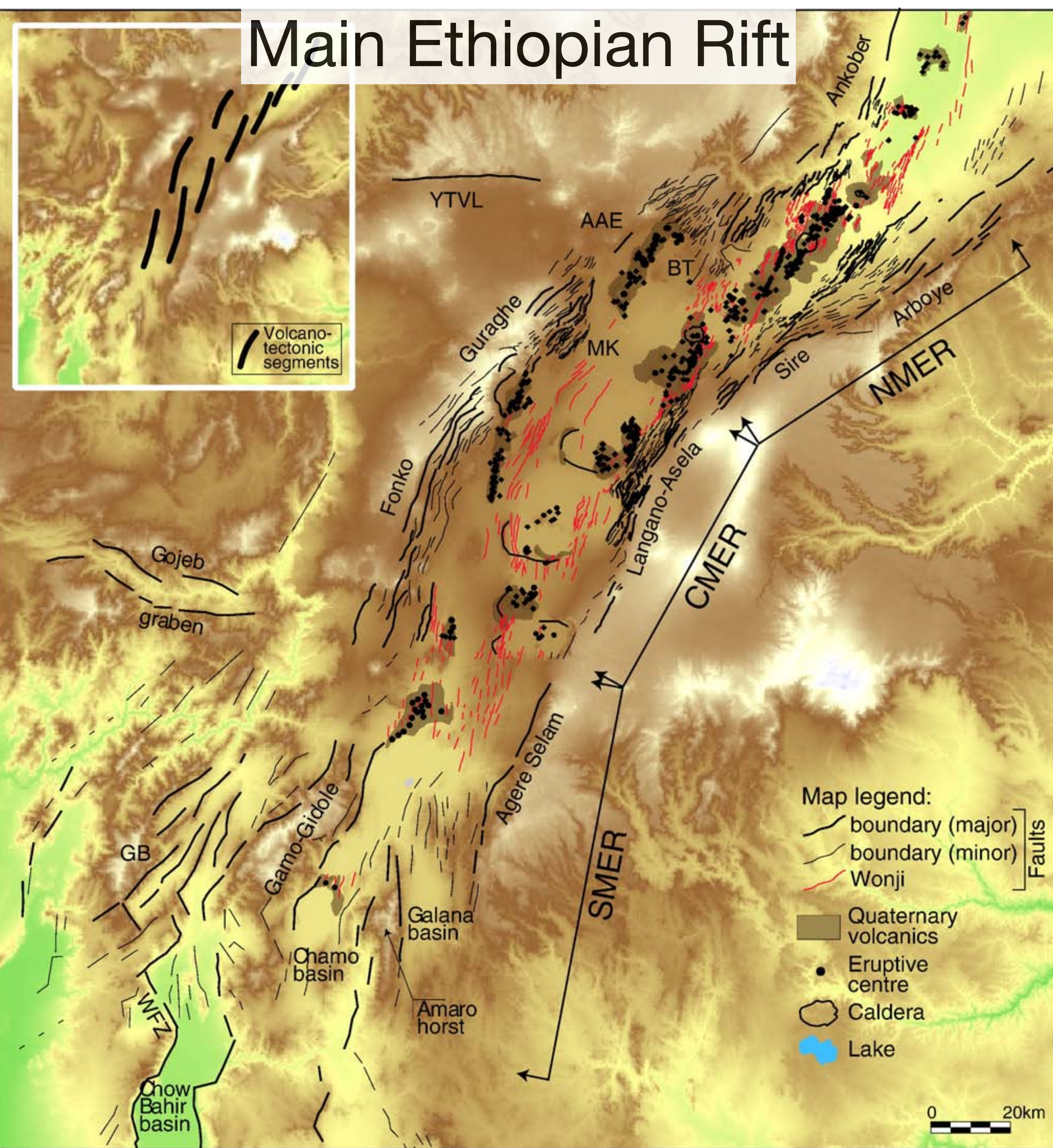
Consequences of preexisting plastic damage variations (“inheritance”)



Influence of orogenic collision inheritance (Mid-Norwegian margin)

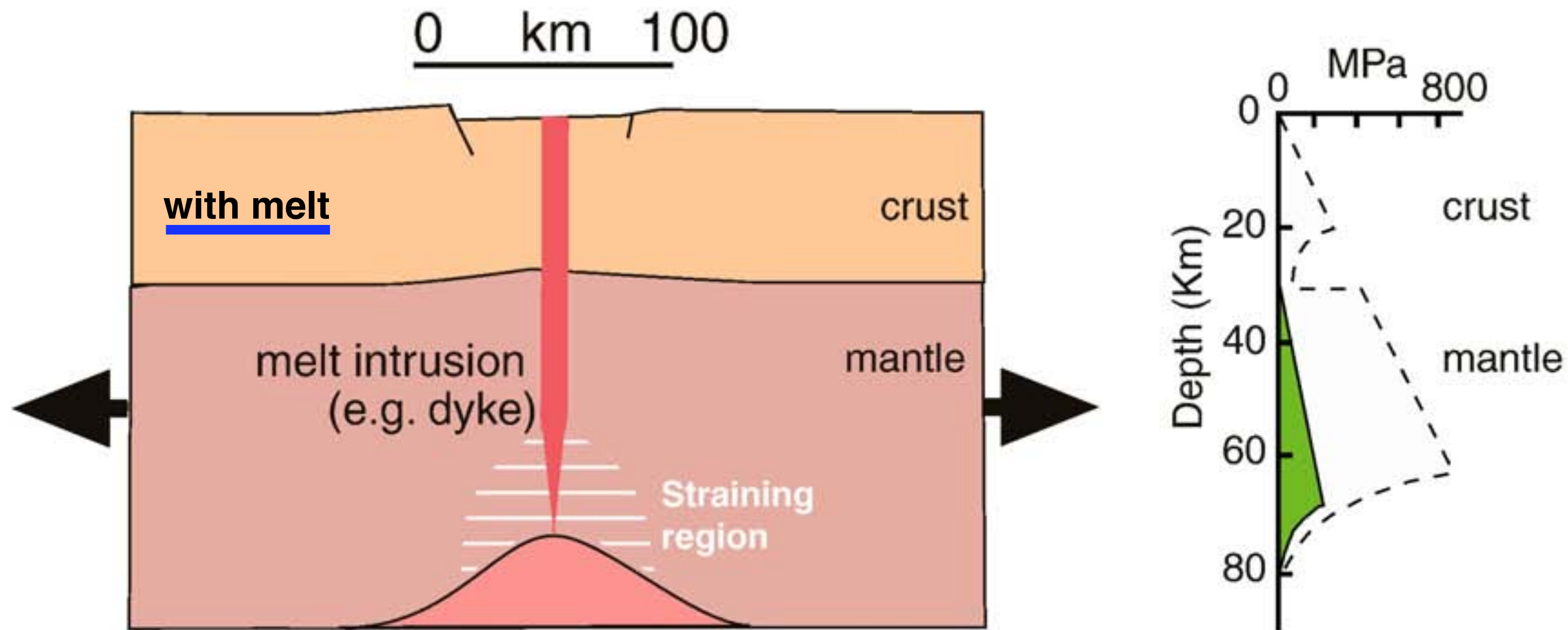


Main Ethiopian Rift



Ebinger 2005, Corti 2009, Boccaletti et al 1998, Casey *et al* 2006

Magma reduces the stress required for rifting

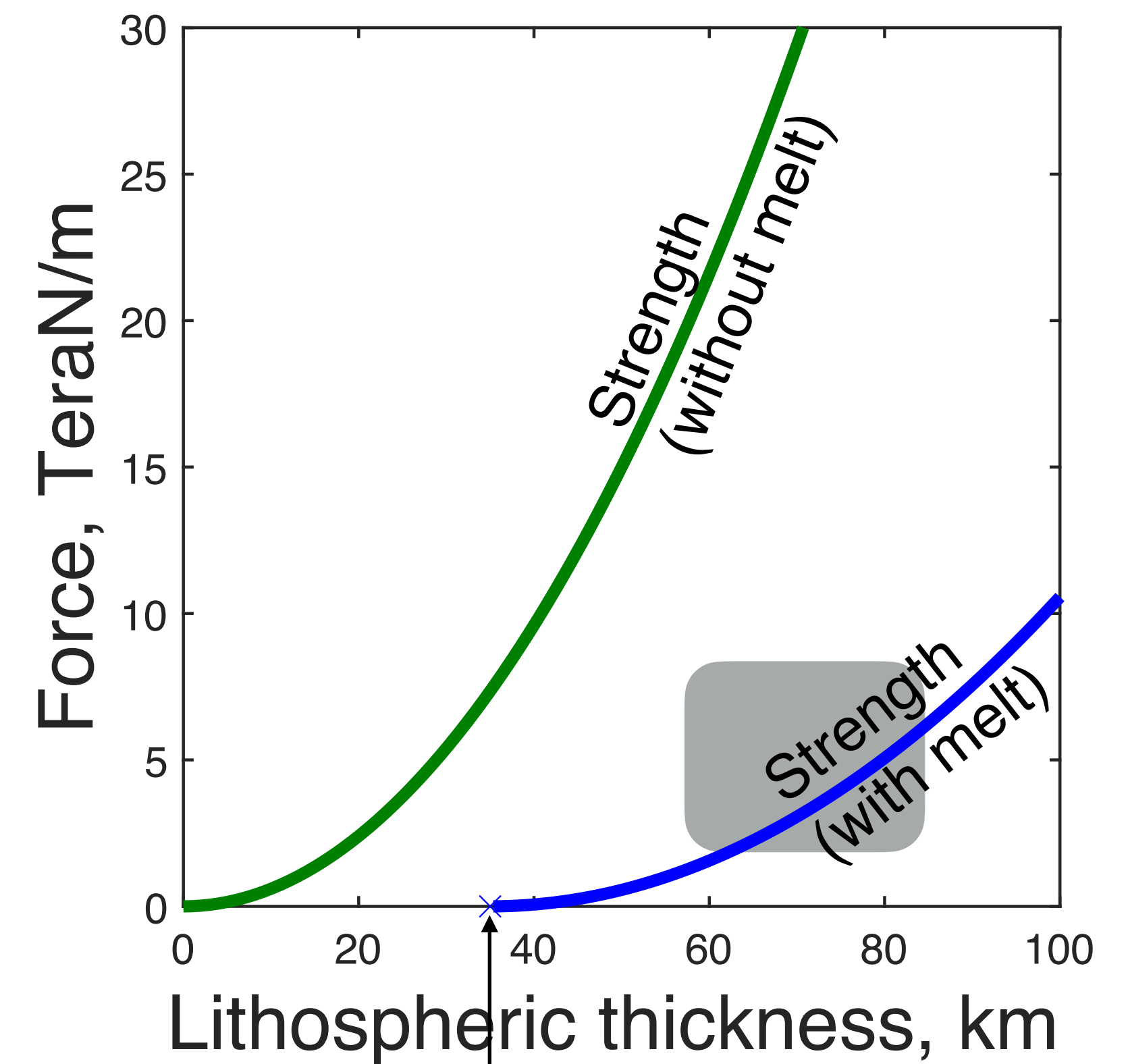
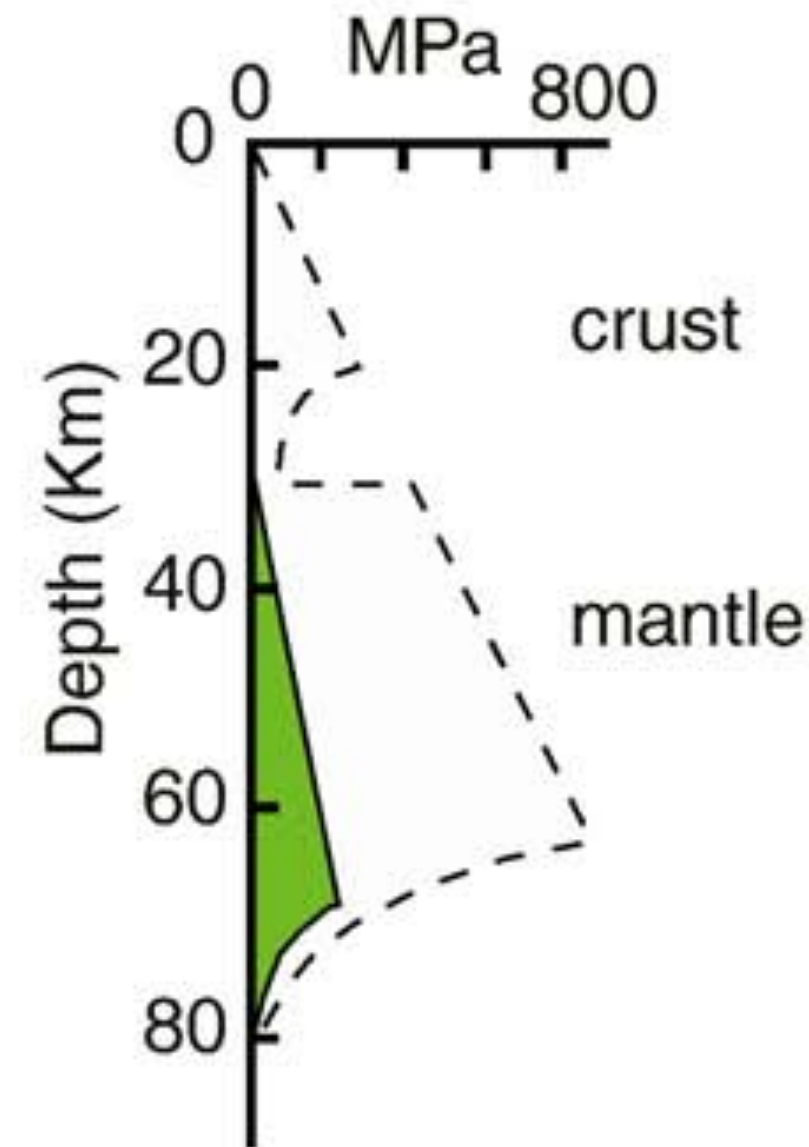
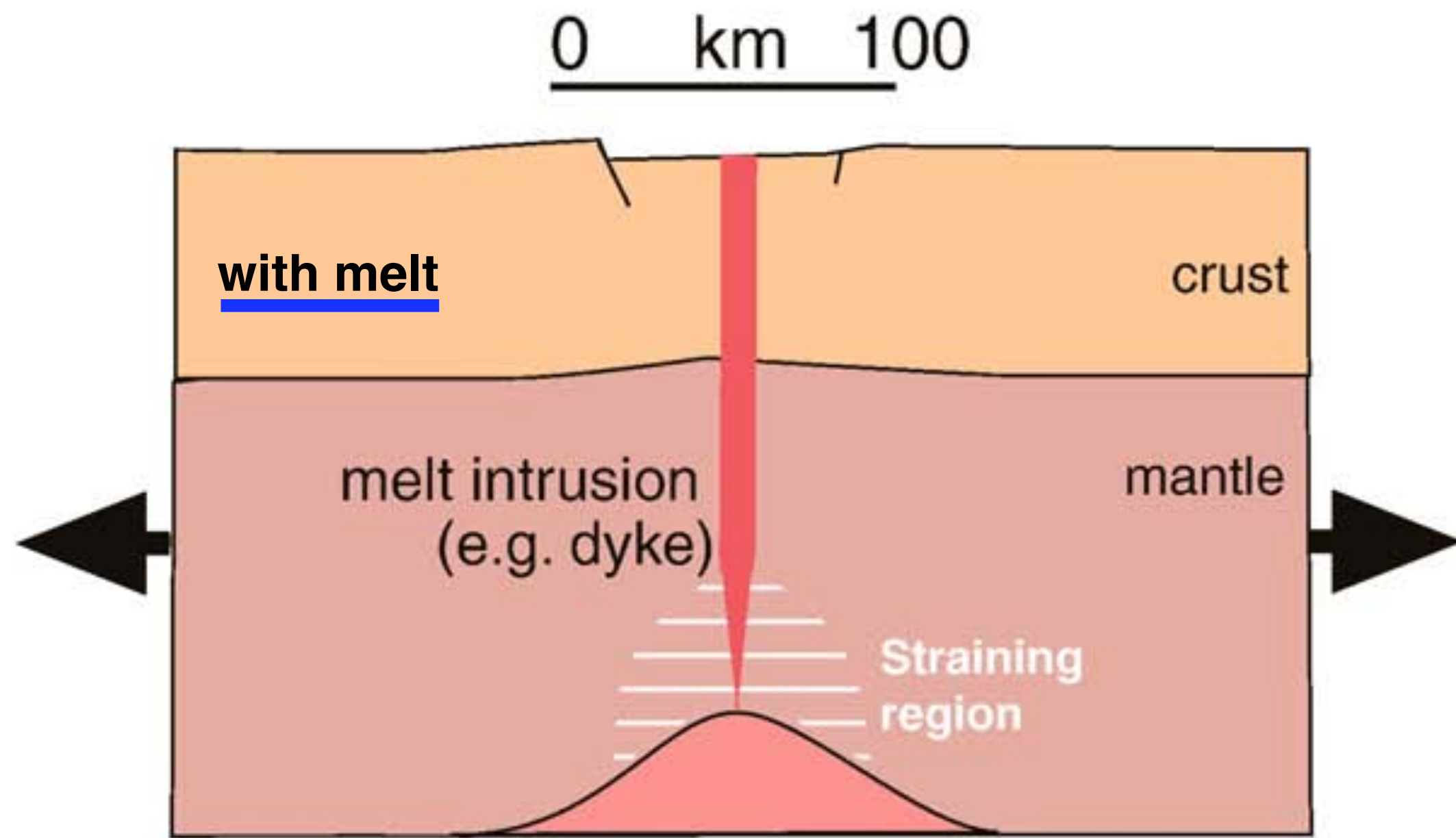


Assume zero cohesion (fracture toughness)

$$F = \Delta GPE = \int_{-L}^0 [\rho_f - \rho_{\text{lith}}(z)] g z \, dz$$

$$\rho_f \approx \rho_c < \rho_m \quad \text{Buoyant magma}$$

Magma reduces the stress required for rifting

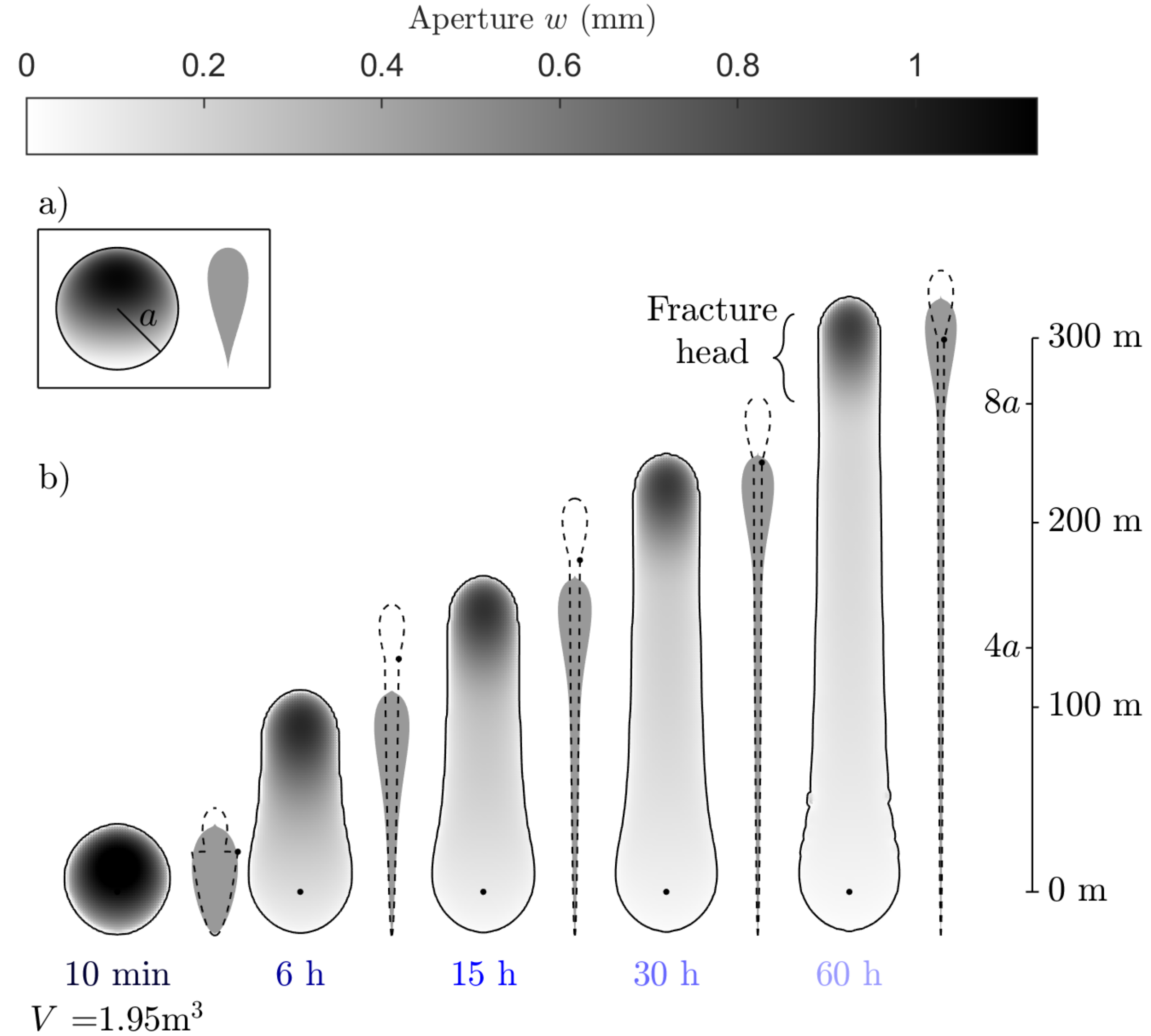


Assume zero cohesion (fracture toughness)

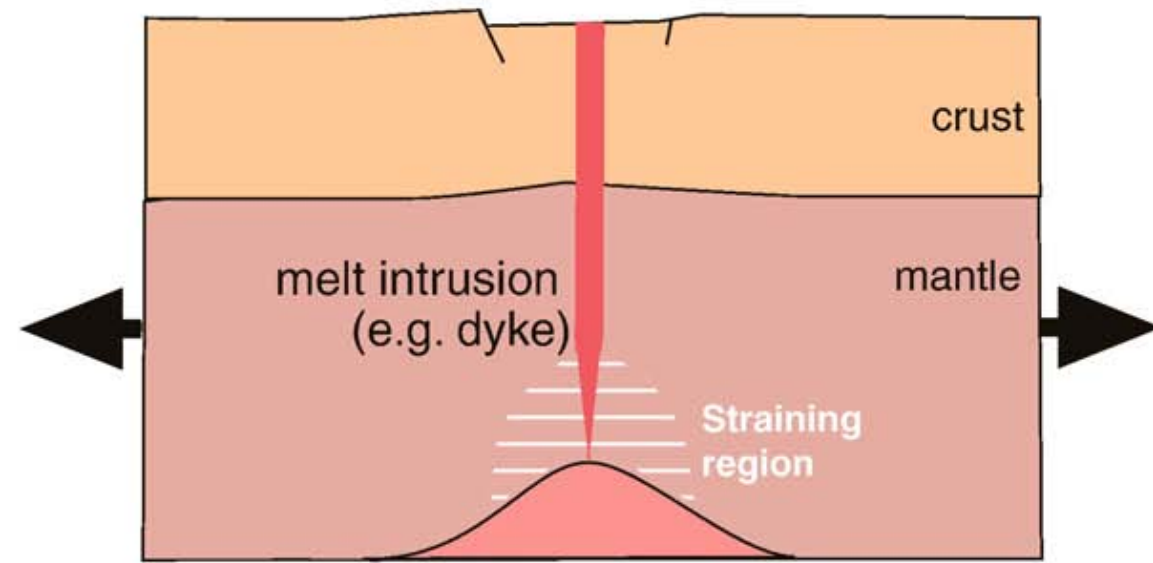
$$F = \Delta GPE = \int_{-L}^0 [\rho_f - \rho_{\text{lith}}(z)]gz \, dz$$

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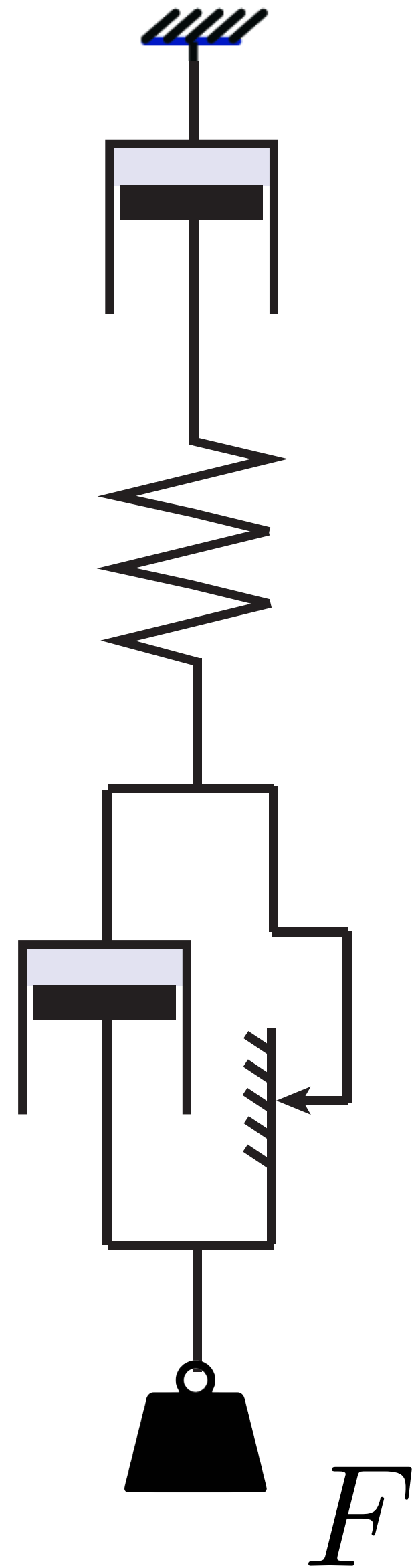
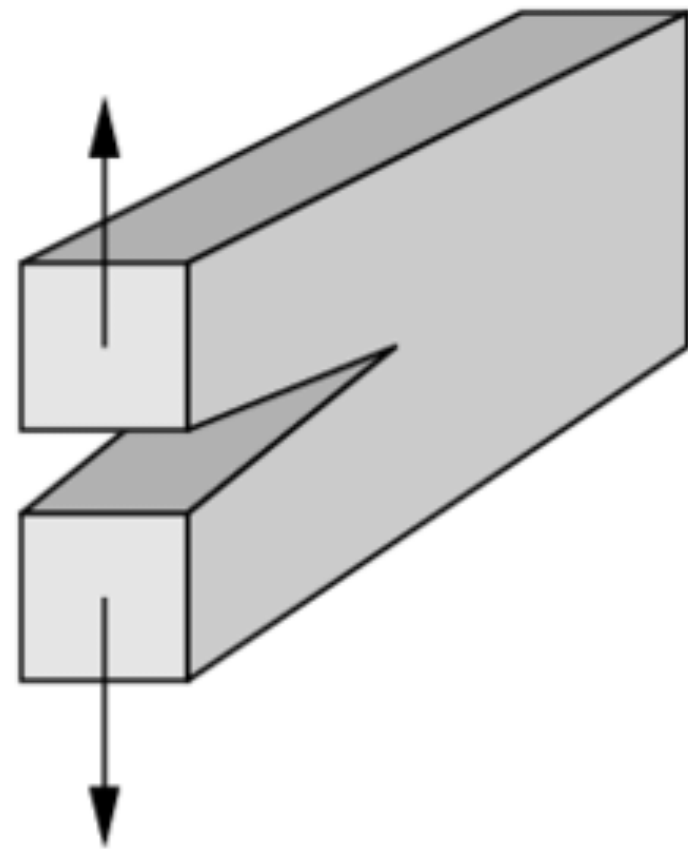
Fluid-driven fracture (e.g., hydrofracture)



Using plasticity to represent dikes?



Mode I



dashpot representing *creep deformation*

extends at a rate $\propto F$

spring representing *elastic deformation*

extends by an amount $\propto F$

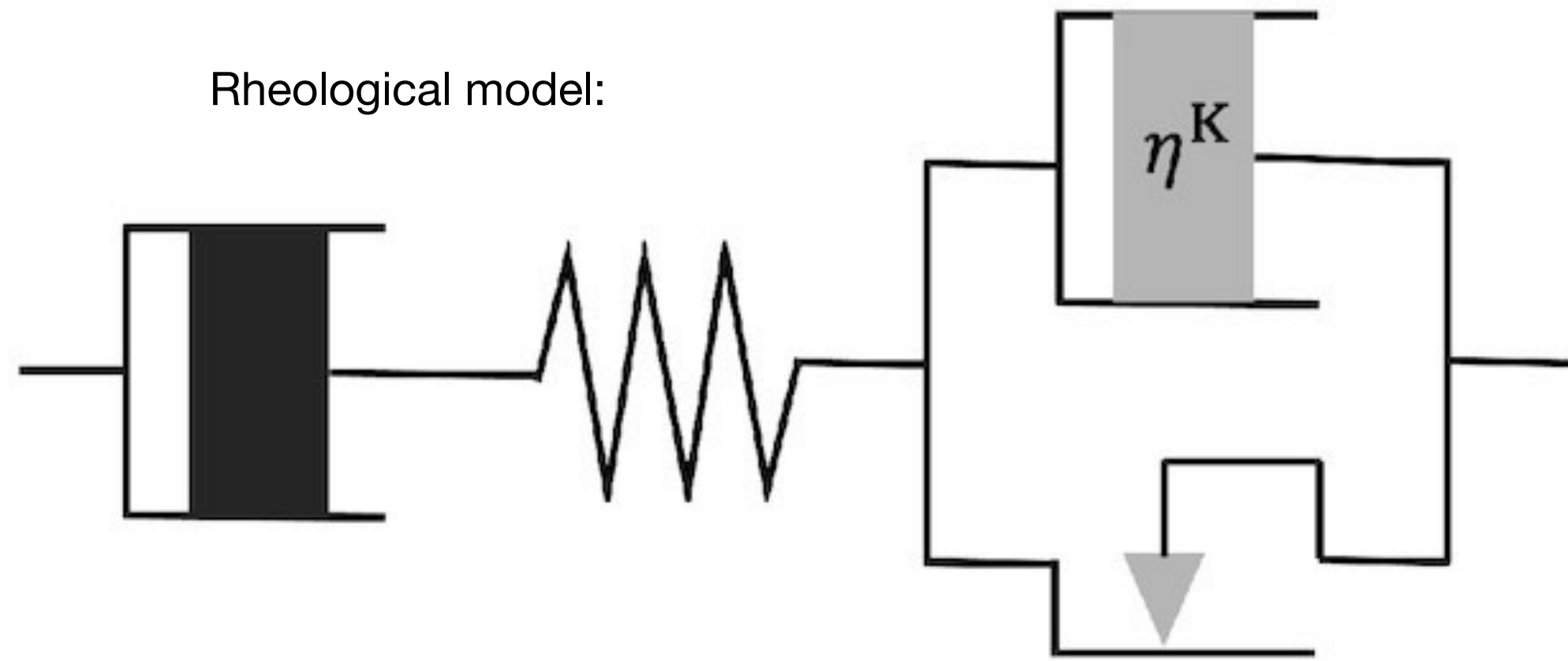
slider representing *fracture and frictional slip*

extends if $F > F_{\text{crit}}$

(at a rate determined by dashpot)

A poro-viscoelastic–viscoplastic model of diking, faulting & rifting

Rheological model:

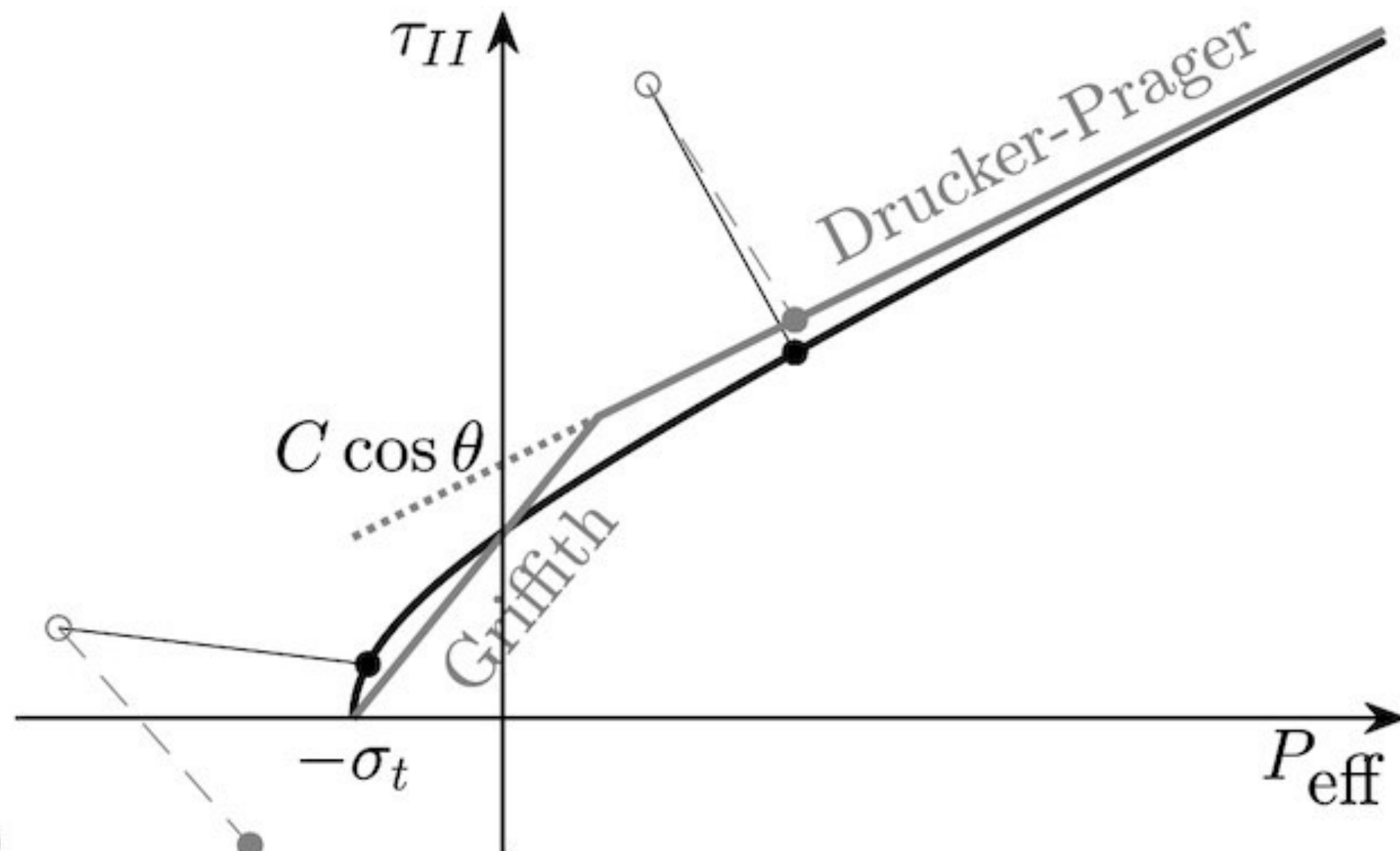


(a) viscous elastic viscoplastic

3 cases:

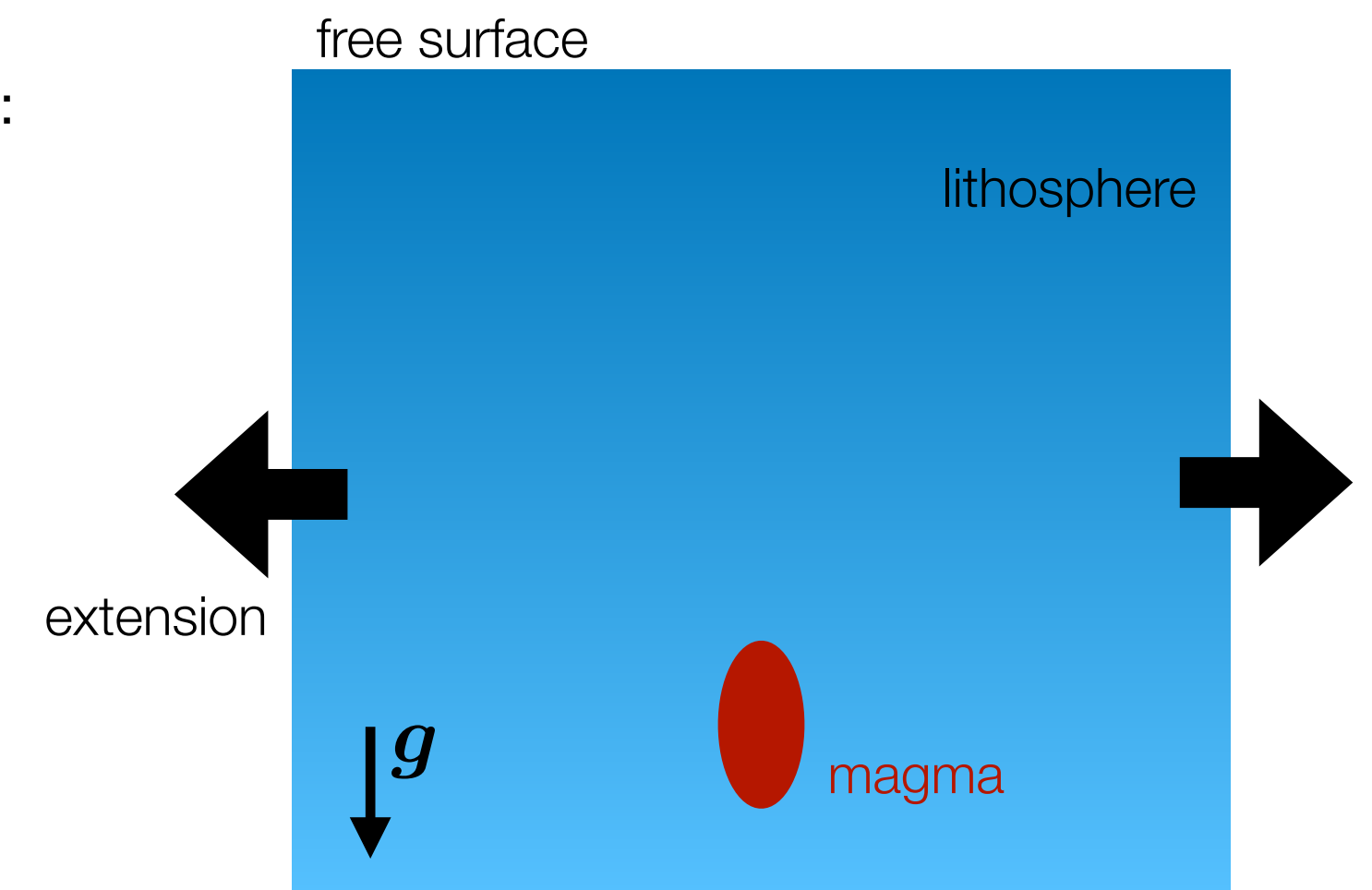
$$\eta_K = \begin{cases} 10^{21} \text{ Pa s} \\ 10^{20} \text{ Pa s} \\ 10^{19} \text{ Pa s} \end{cases}$$

Plastic failure envelope:

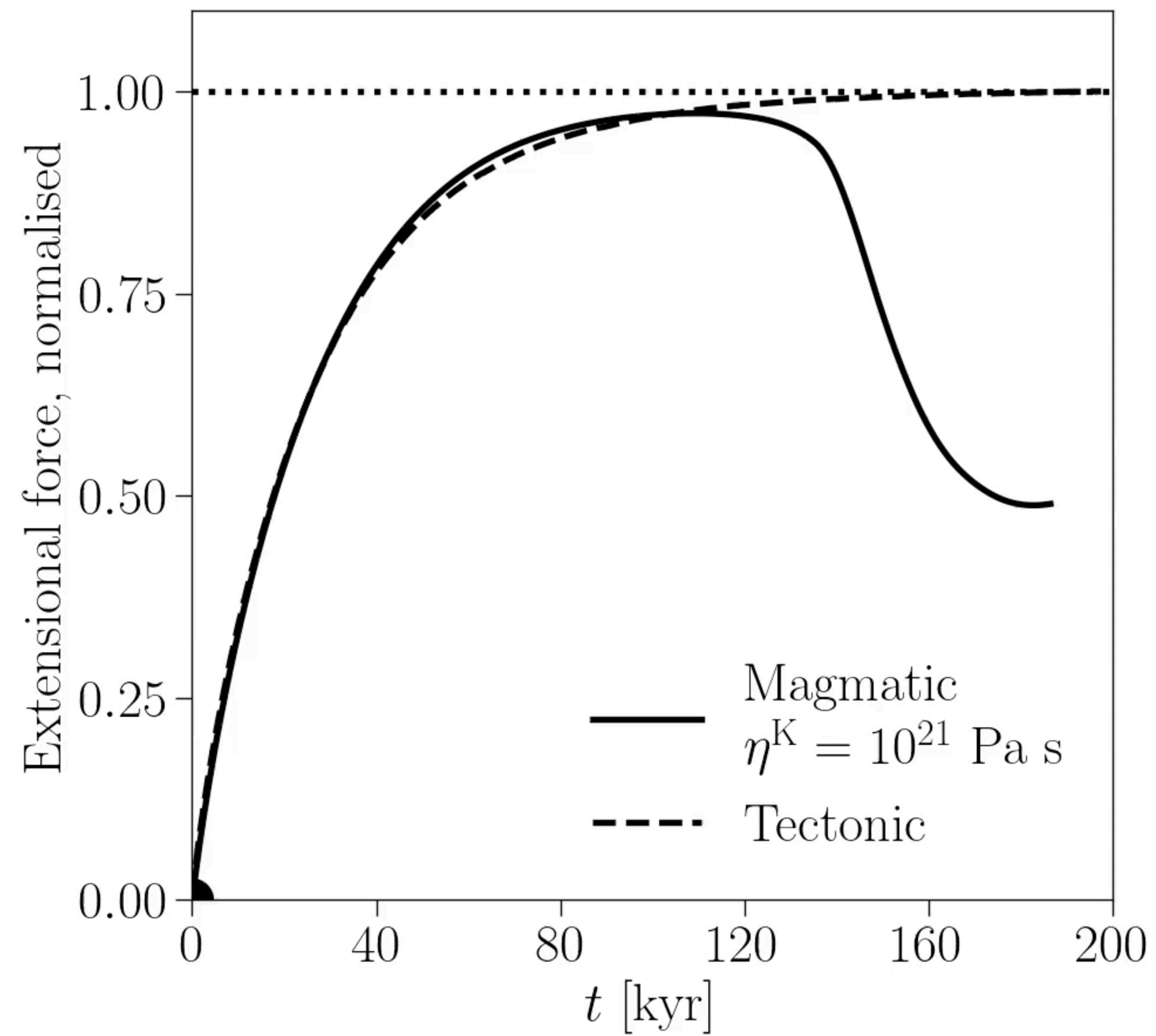
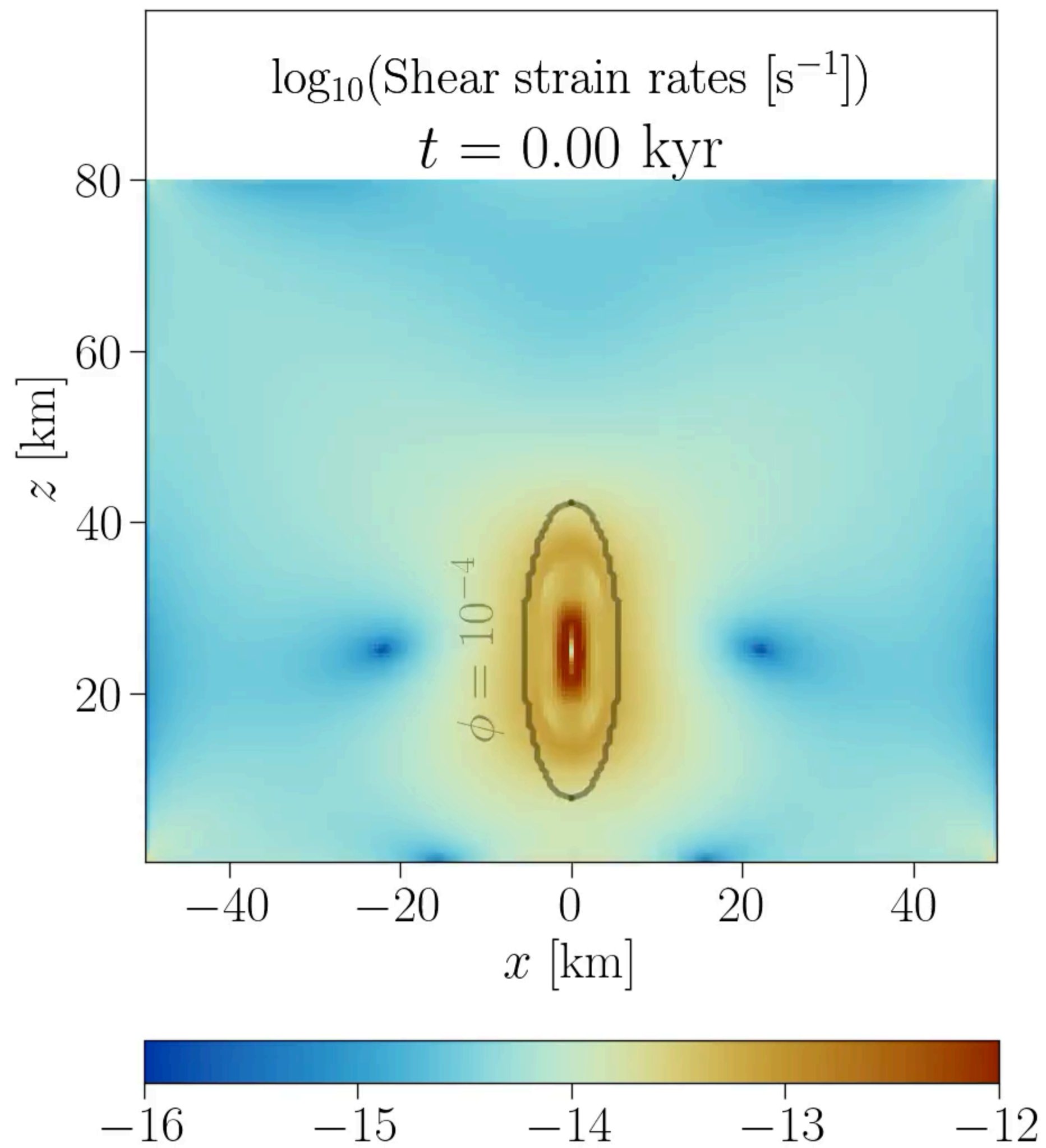


(b)

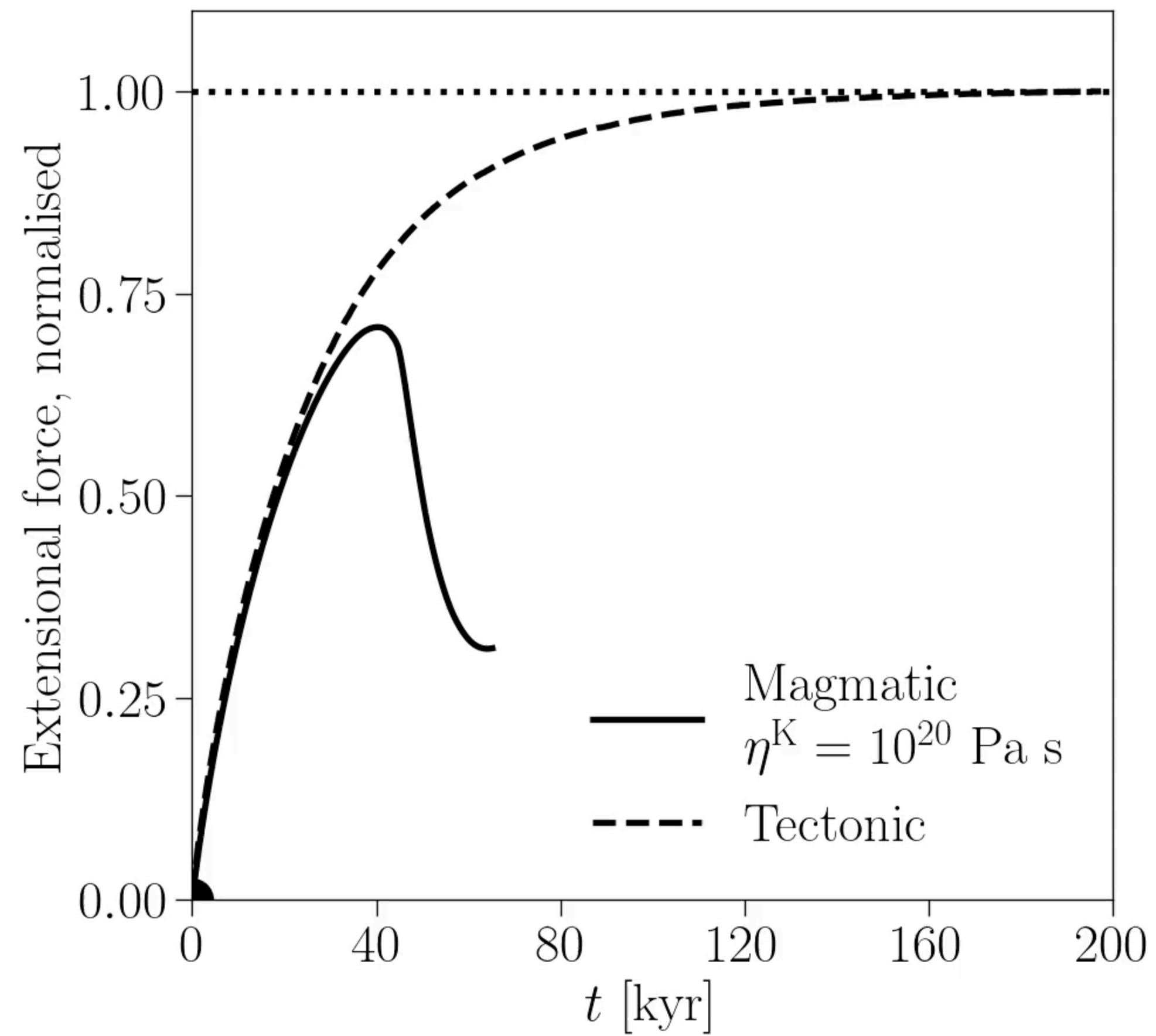
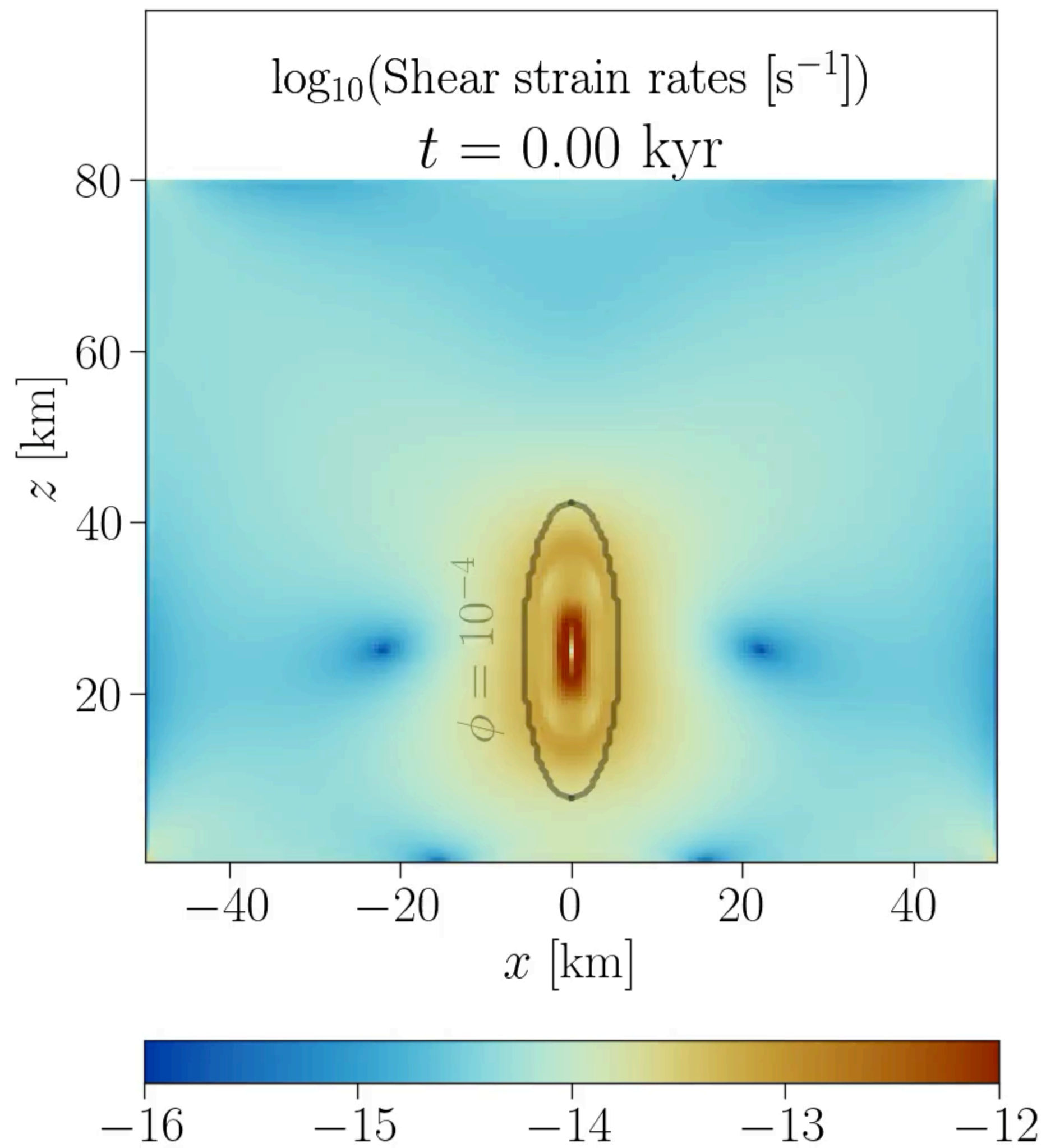
Model set-up:



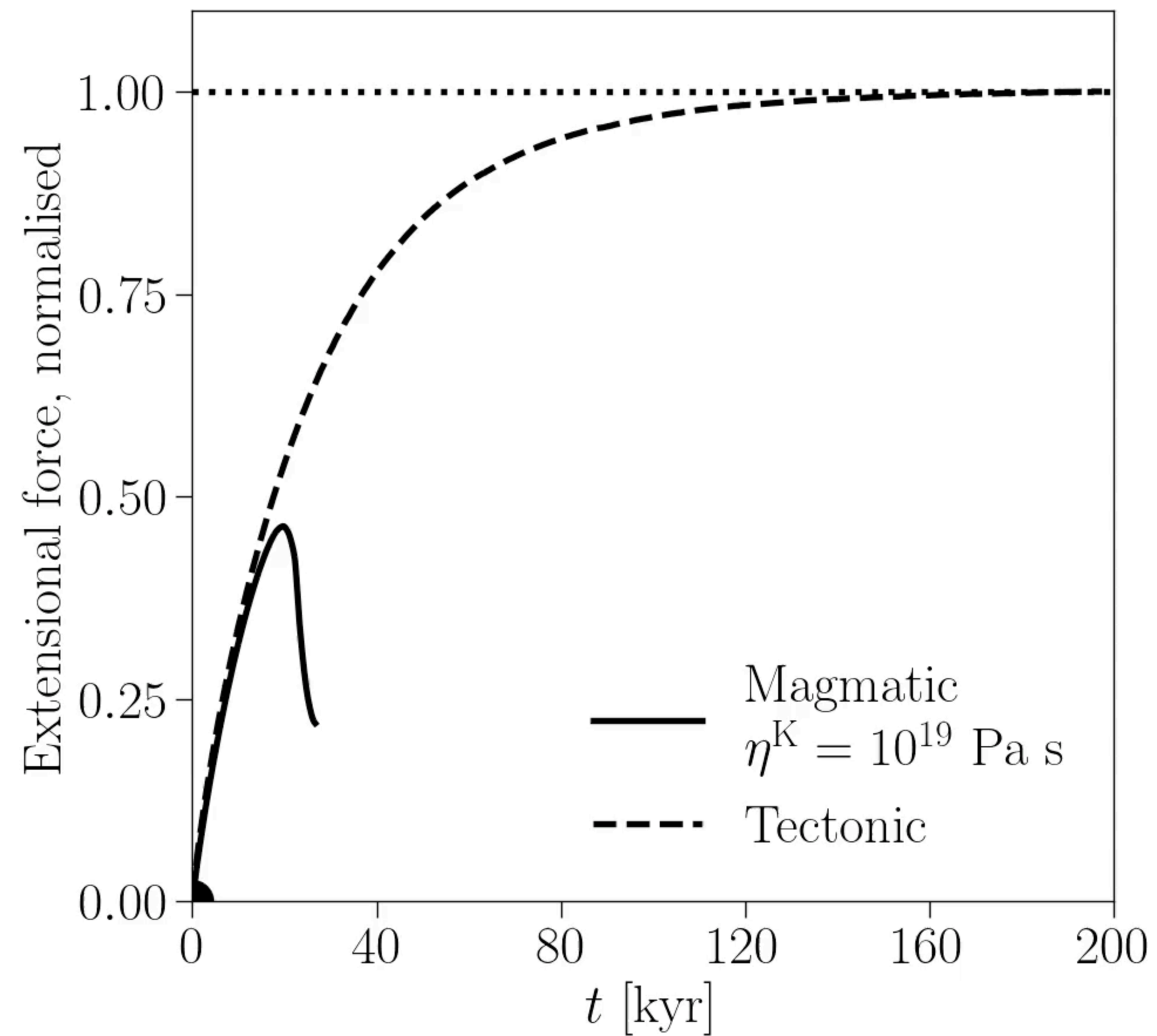
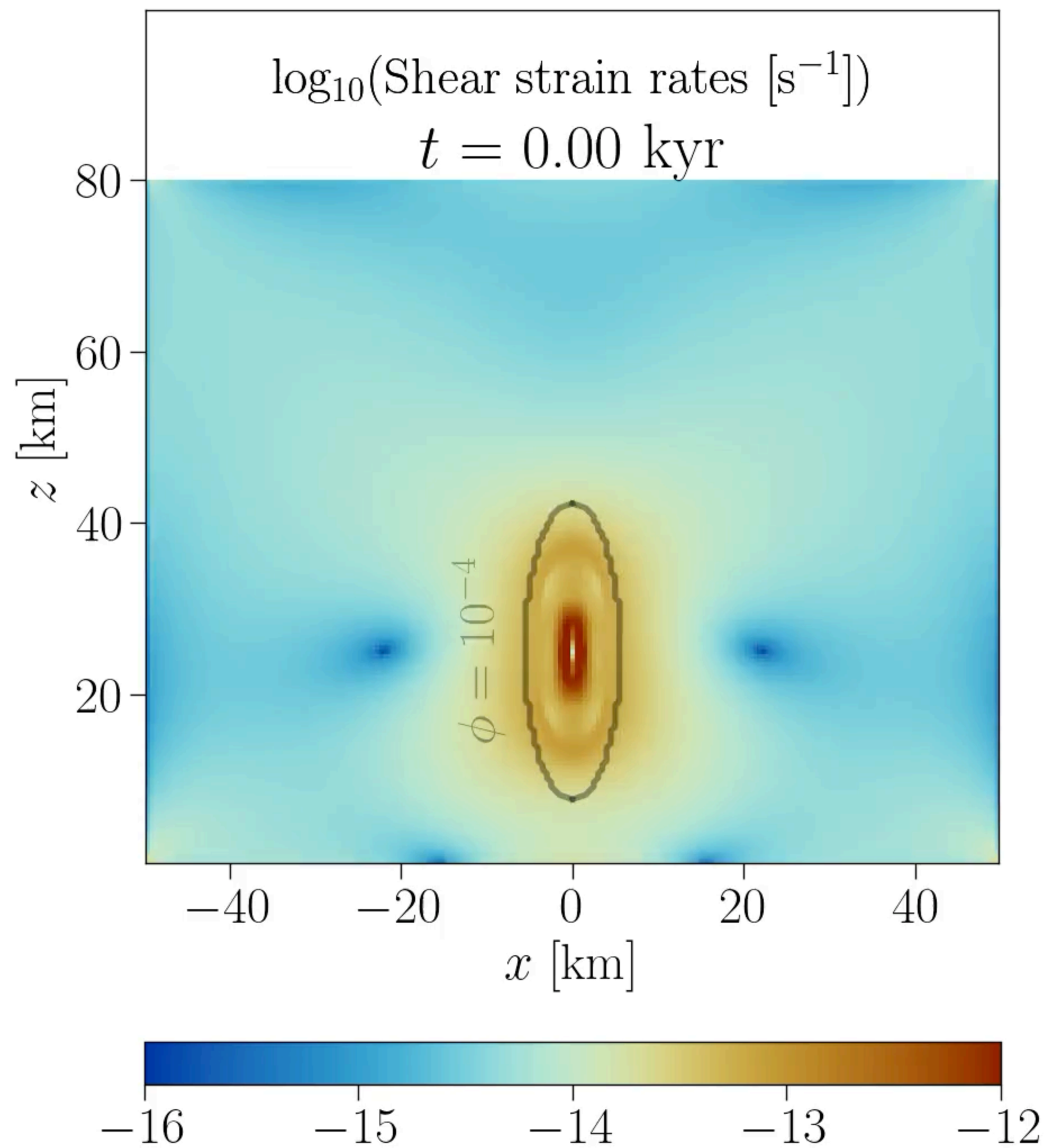
A poro-viscoelastic–viscoplastic model of diking, faulting & rifting



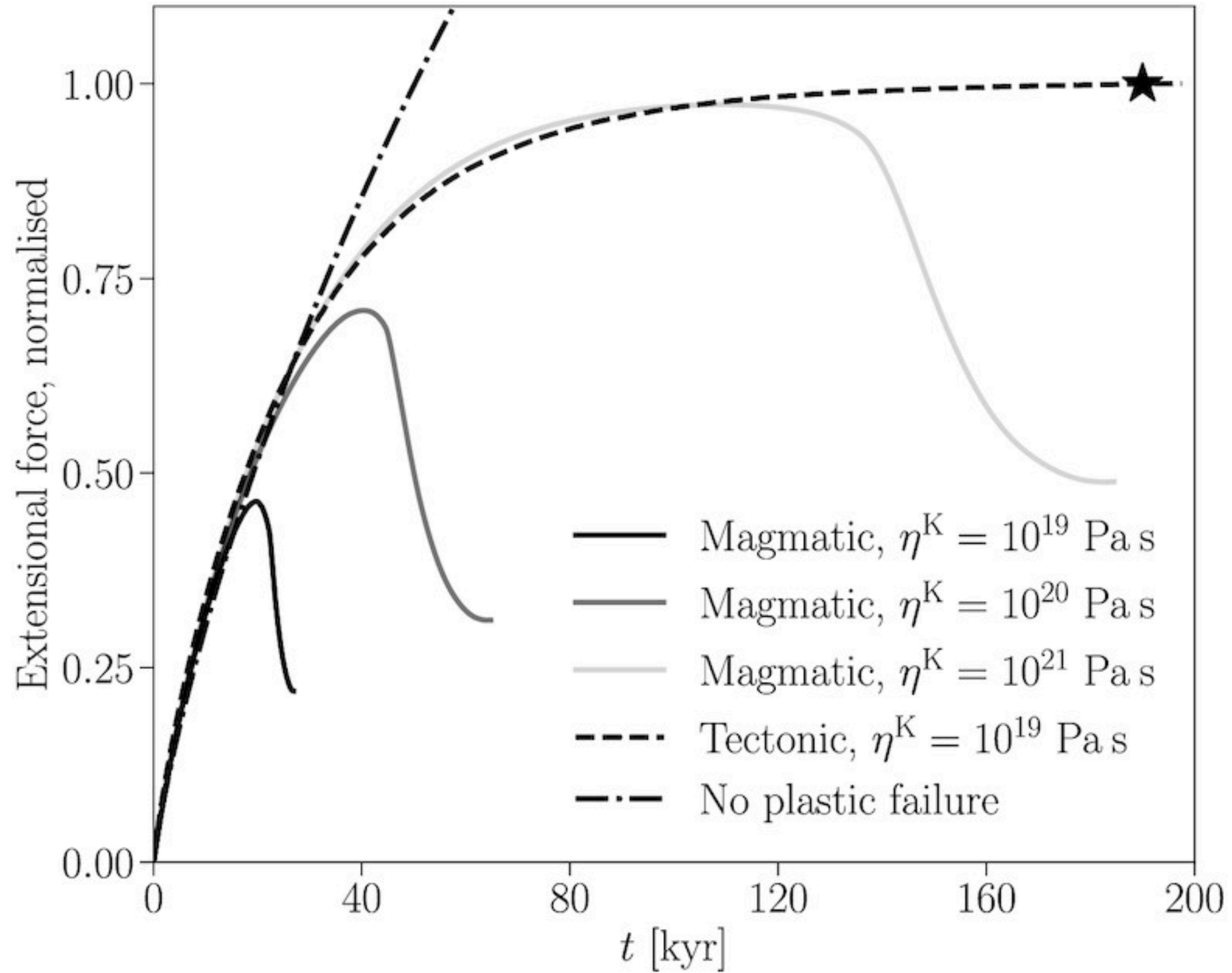
A poro-viscoelastic–viscoplastic model of diking, faulting & rifting



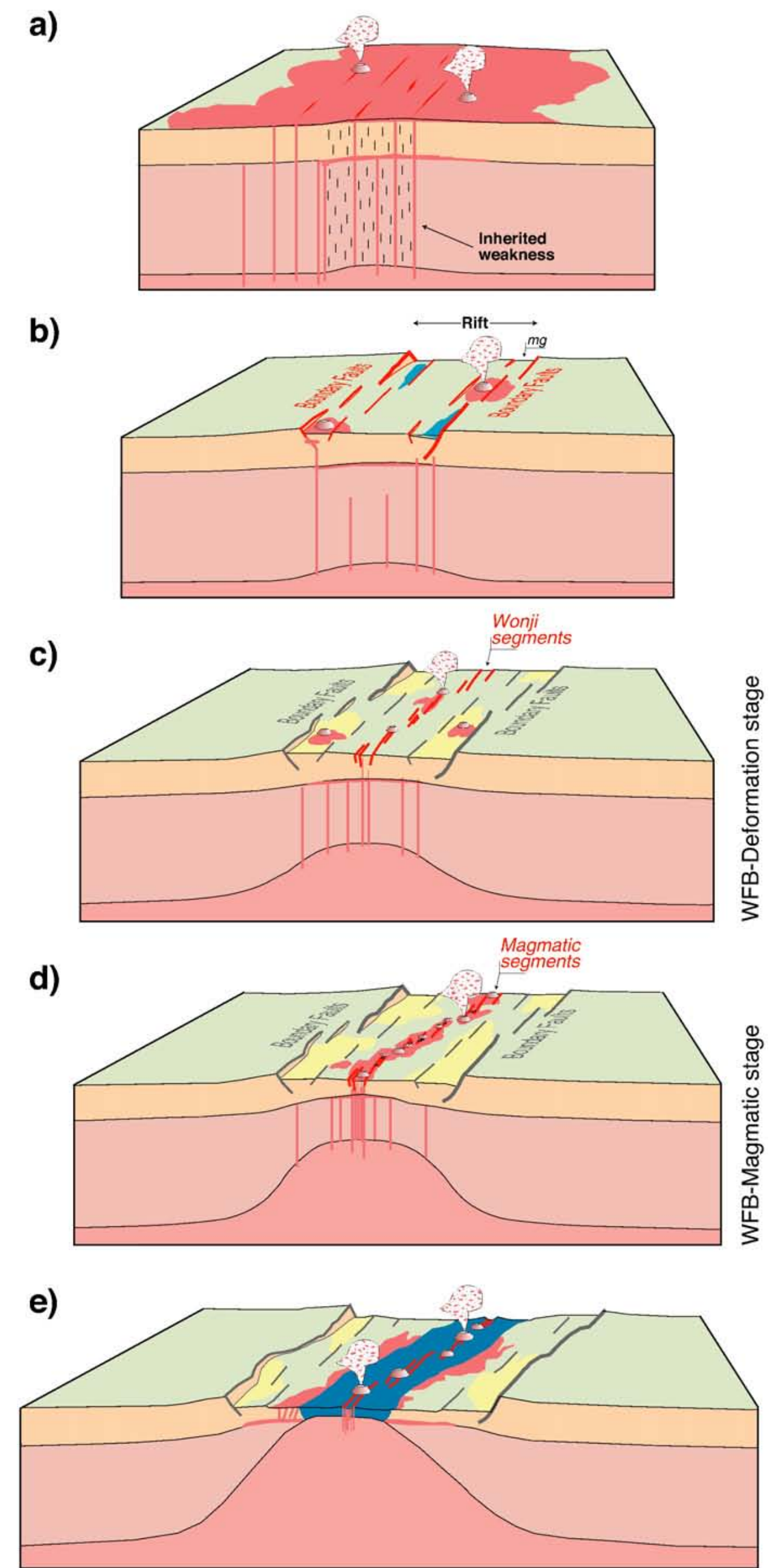
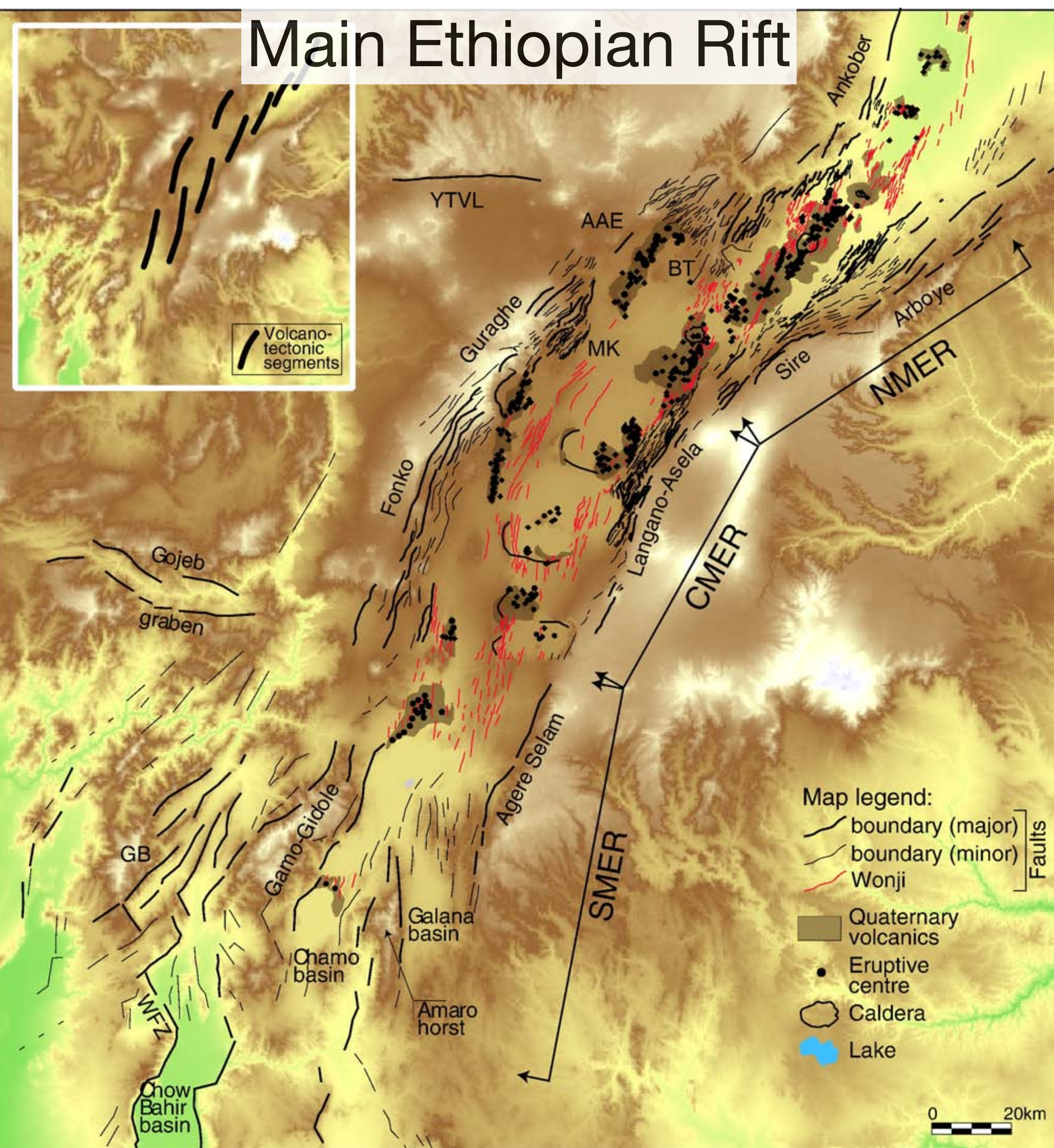
A poro-viscoelastic–viscoplastic model of diking, faulting & rifting



The force for magmatic rifting



Main Ethiopian Rift



Ebinger 2005, Corti 2009, Boccaletti et al 1998, Casey *et al* 2006

Summary and questions

- Earth is a *thermal engine* that converts heat (radiogenic, primordial) into work by *convection*.
- Plate tectonics is the surface expression of mantle convection.
- Continents rift under the convective forces of slab pull, topographic/GPE gradients, and basal drag.
- Tectonic forces are of order 10 TN/m.
- Why do rocks yield and what do they remember of past yielding? What is mechanical role of magma at rifts and mid-ocean ridges? How do faults and dikes interact mechanically?

