Modeling the Impact of Climate Change on Flood and Drought: Case Study Awash River Basin, Ethiopia



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Climate?

Is the weather in some location/place averaged over some long period of time (Long-term average of weather condition)

Climate Change?

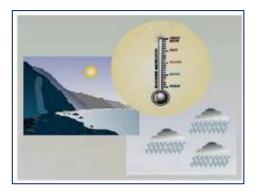
The change in climate variable that occurs over a longer period of time typically over decades and centuries and may not return to it former state unless significant measures are taken place.

Climate Variability?

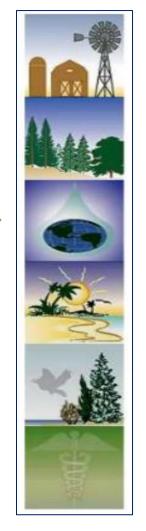
The change that occurs within smaller/shorter time frame such as a month, a season or a year and may return to its former state.



Consequence of Climate Change



Sea level rise
 Change in Temperature
 Change in Precipitation
 Extreme events (flood and drought)



Agriculture and food security (crop yields, irrigation demand....)

Forest

(composition, health and productivity....)

Water resources

(water supply, water quality,....)

Coastal areas

(erosion, inundation, cost of prevention,...)

Species and natural areas

(biodiversity, modification of ecosystems....)

Human health

(infectious diseases, human settlements,....)



Increased surface

concentration in the

Unexpected hazards

such as drought,

flooding, global

warming, rise of

ocean and lake

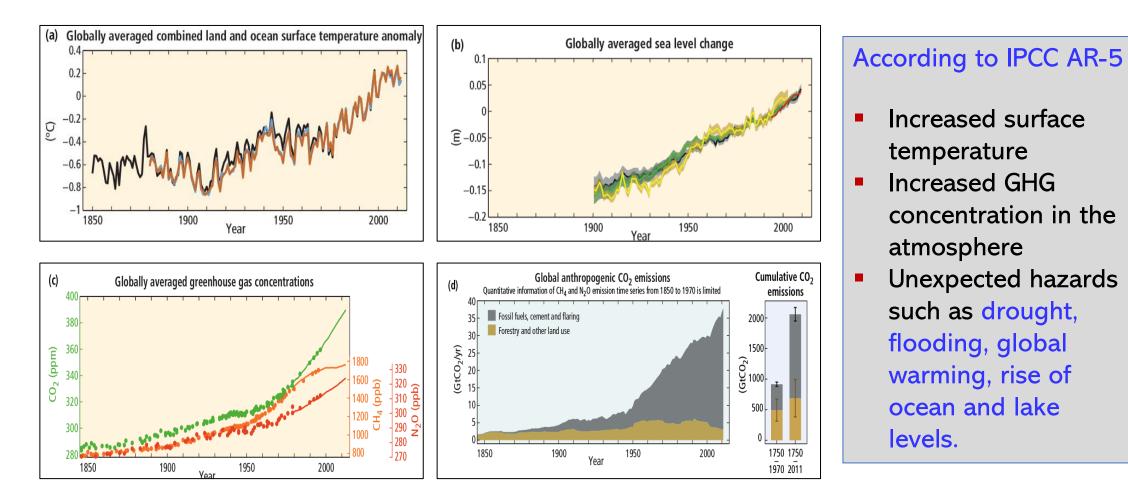
levels.

temperature

atmosphere

Increased GHG

Basic Observation of Climate Change in IPCC 5th Assessment Report





Basic Observation of Climate Change in IPCC 5th Assessment Report

Temperature Change

- The global average combined land and ocean surface temperature shows a warming of 0.85 (0.65 to 1.06) °C.
- Each of the last three decades has been successively warmer at the Earth's surface than any preceding decade since 1850.

Greenhouse Gas Emission

Anthropogenic greenhouse gas (carbon dioxide, methane, and nitrous oxide) emissions have increased since the preindustrial era, driven largely by economic and population growth, and are now higher than ever.

Sea Level Rise

Over the period 1901 to 2010, the global mean sea level rose by 0.19 (0.17 to 0.21) m.
 The rate of sea level rise since the mid-19th century has been larger than the mean rate during the previous two millennia.



Climate change could intensify hydrological extremes, changing not just the magnitude but also the timing of flood and drought events (Lavers et al., 2015).

Ethiopia

- ✓ High potential for extreme events (floods and droughts), wet seasons can be wetter and warm seasons can also be warmer, (NAPA, 2007).
- Ethiopia is subject to high climate variability and change (IPCC, 2007), experiencing frequent floods and droughts, particularly in the Awash River Basin.
- The current scenario indicates that flood and drought are the recurrent common phenomena of Awash Basin, with devastating effects on environmental, social, and economic loss.

1.1. Objective

→ To assess the impacts of climate change on hydrological extreme events (flood and drought) under current and future climate change scenarios

Where is our research site?

- Conducted in the middle and lower Awash River Basin, Ethiopia
- Geographical location, north-eastern part of Ethiopia,

8°49' to 14°30'N latitude and 39°34' to 42°28'E longitude.

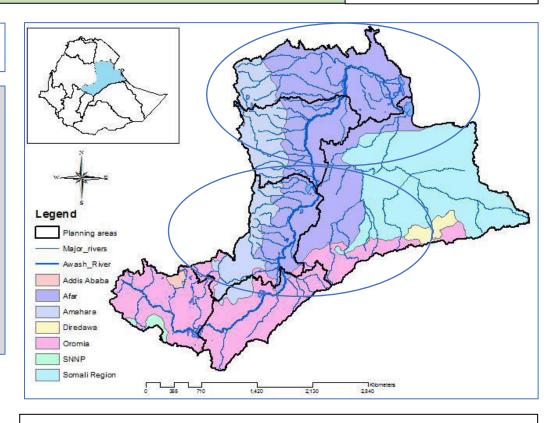
Arid and semi-arid climate with low and erratic rainfall.

Middle Awash Valley

- \rightarrow Between Awash station and the Mille river.
- \rightarrow The altitude varies from 500m to1000m a.s.l.

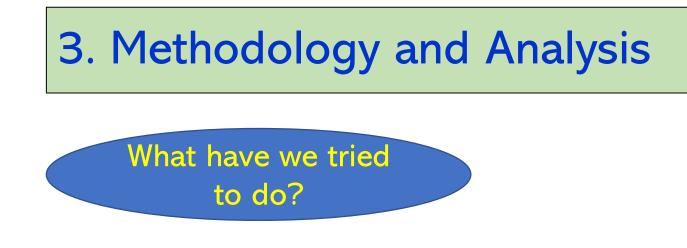
Lower Awash Valley

→ An altitude between 250m and 500 m with the mean annual rain fall of less than 200mm.





2. Study area



 This study mainly focused on assessing the potential impacts of climate change on hydrological extremes (flood and drought) with the application of

 \rightarrow a physically based hydrological model HEC-HMS,

→Standardized Precipitation Indices (SPI) and

 \rightarrow RCM model output for future climate projection in the basin.



Data Collection

- 1. Climate Data
 - \rightarrow Observed/station data
 - → Three RCM model output (MIROC5-RCA4, CSIRO-RCA4, and CNRM-RCA4)
- 2. Hydrological Data -----River flow data
- 3. Spatial Data -----DEM, Digital Stream Network, Soil map, and LULC



Table. Summary of input dataset, including data-type and their sources.

Data	Source	Description		
Climate				
Observed data RCM model data	Ethiopia national meteorology agency CORDEX Africa	Daily maximum and minimum temperatures and precipitation (historical and future period)		
Hydrological data	Minster of Water Resource (MoWR)	Daily streamflow (1996–2015)		
Soil map	Digital Soil Map of the World (DSMW)	Soil classification and properties		
Land use	ESA-CCI land use (<u>http://esa-landcover-cci.org/</u>)	Land use classification – cropland, forest, etc.		
DEM	ASTER (<u>https://earthdata.nasa.gov/</u>)	Elevation, overland, channel slopes, boundary		

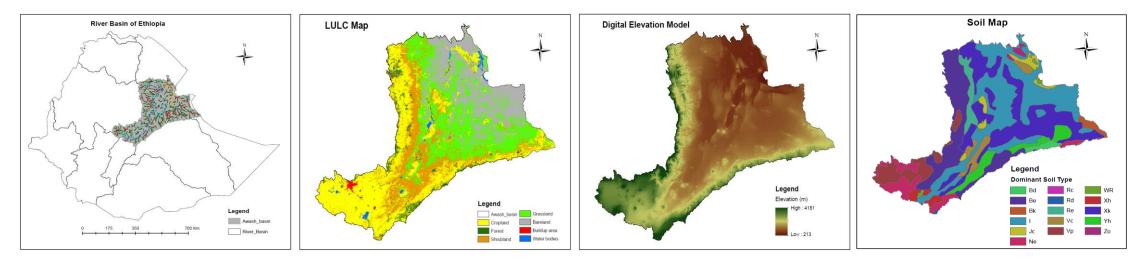


Figure. Input spatial data for HEC-HMS model for the Awash River Basin



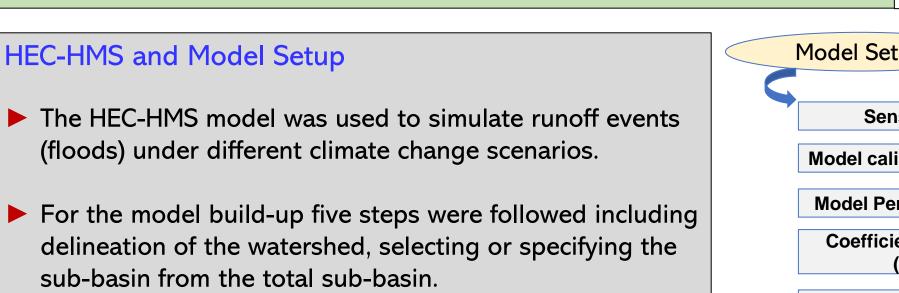
Climate Change Scenario and RCM output

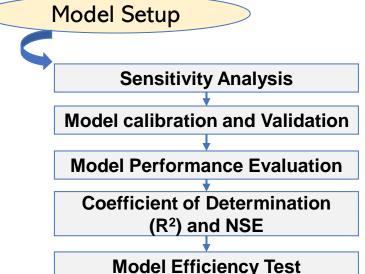
- The future climate variables were projected using the RCM model based on two RCP scenarios.
 - \rightarrow RCM models: MIROC5-RCA4, CSIRO-RCA4, and CNRM-RCA4 \rightarrow Scenario Generation: RCP scenarios (RCP4.5 and RCP8.5)

Bias Correction Method

 Linear-scaling approach: due to its suitability and simplicity for bias correction on a daily basis of precipitation and temperature data.

□ The future scenarios were developed by dividing the future time series into two periods: \rightarrow 2030s (2011-2050) and 2090s (2051-2100).





University of Tsukuba

Model calibrations and validation

3. Methodology and Analysis

The model's performance was evaluated through the calibration and validation process and then the calibrated model was then applied to various climate scenarios.

- The model ran for 19 years (1996-2015) using available observed data.
- Warm-up period (1996-1998) for normalization of the model.
- The model calibration (1999-2011) and model validation (2012-2015).

 Table. Classification of drought

SPI value	Drought category
2.0+	Extremely wet
1.5 to 1.99	Very wet
1.0 to 1.49	Moderately wet
-0.99 to 0.99	Near normal
-1.0 to -1.49	Moderately dry
-1.5 to -1.99	Severely dry
-2 and less	Extremely dry

events based on SPI value (McKee et al., 1993)

Drought Analysis Based on Drought Indices

- The standardized precipitation index (SPI) was used to assess the meteorological drought in terms of severity and duration.
- The SPI is the most widely used drought index which is recommended by organizations like WMO and NOAA.
- The SPI is estimated based on long-term precipitation records that are fitted to a probability distribution
- SPI uses only precipitation data for its calculation.

4. Results and Discussion



Model Calibration and Validation Results

- The performance of the HEC-HMS model was evaluated through a calibration and validation process using observed streamflow data over a period of 1996-2015 (Sisay et al., 2017).
- The model performed well for our study area with reasonable accuracy as shown in Table both Nash-Sutcliffe Efficiency (NSE) and coefficient of determination (R²) > 0.7 for both the calibration and validation periods (1999-2011 and 2015-2015).

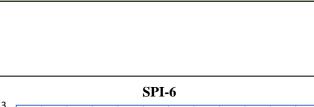
Table. Summary of model performance criteria for calibration and validation

	Middle Awash		Lower Awash	
Period	R2	NSE	R2	NSE
Calibration (1999-2011)	0.78	0.8	0.86	0.79
Validation (2012-2015)	0.82	0.73	0.83	0.71

Drought Analysis Results The drought indices for the 6and 12-month time scales

- (SPI-6 and SPI-12) were selected for the study of drought characteristics.
 A drought event occurred when the drought index (SPI)
- when the drought index (SPI) values were less than -1.0 (Table).
- The overall categories of drought are considered (mild drought and above), areas in the Middle and Lower Awash Basin are most frequently hit by droughts.





Time (in Years)

SPI-6 Value

2031 2041 2051

Years

2061

2071

2

SPI value

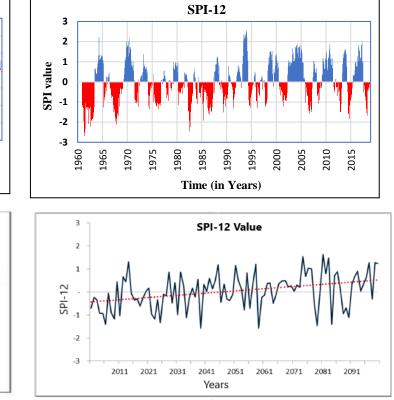
-3

SPI-6

2011

2021

1960 1975 1975 1975 1985 1985 1990 1995 1995 2000 2000 2010 2015 2015 2015





4. Results and discussion

Flood Analysis Results

Table. changes in flood magnitudes between the current; and the 2030s and 2090s time periods.

Return		5	10	25	50	100	
periods		Percentage (%)					
Periods	2030s	8.7	4.2	1.5	-2.6	-5.5	
	2090s	10.5	8.9	8	7.3	7.2	



Figure. Flood frequency curve for Awash River based on future scenario simulation.

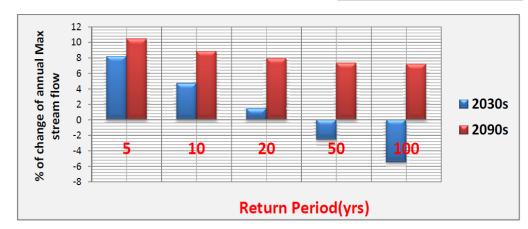


Figure. Changes in flood magnitudes between the current and the 2030s, 2090s the time period corresponding to the return period

- Figure and table, present the percentage increase in peak flow of river corresponding to different return periods (5, 10, 25, 50, 100) by the 2030s and 2090s.
 These results indicate an average increase in flood
 - events.
- Moreover, the analysis indicates that the overall increasing trend in the frequency of flood events is not linear.



5. Conclusion



- → This study addresses climate change impact assessment on the hydrological extremes (drought and flood)
- \rightarrow For hydrological simulation, we used the HEC-HMS model
 - calibrated and validated in daily and monthly time steps.
 - the model performed well with reasonable accuracy (NSE and R2>0.7).
- → Moreover, drought analysis indicates that the overall categories of drought are considered (mild drought and above), and the basin is most frequently hit by droughts.
- → Furthermore, the maximum river flows in the future will be higher and more variable in terms of magnitude, and irregular occurrence, than at present. It is observed that climate change has a significant impact on the high flow condition of the Awash River.

To investigate the robustness of the results, further research on future extreme events is required by considering multi-model ensemble climate scenarios and socio-economic changes (land use land cover change).

References

- De Luis, M., Gonzalez-Hidalgo, J.C., Brunetti, M. and Longares, L.A. (2011). Precipitation concentration changes in Spain 1946–2005. Nat. Hazards Earth Syst. Sci. 11,1259–1265.
- Fay, M., Block, R.I. and Ebinger, J. (2010). Adapting to Climate Change in Eastern Europe and Central Asia, Vol. 52862. World Bank Group, 180.
- Hare, W. (2003). Assessment of Knowledge on Impacts of Climate Change, Contribution to the Specification of Art, 2 of the UNF CCC. WBGU
- Lavers, D. A., Ralph, F. M., Waliser, D. E., Gershunov, A., and Dettinger, M. D. (2015). Climate change intensification of horizontal water vapor transport in CMIP5. Geophys. Res. Lett. 42, 5617–5625.
- McKee, T.B., Doesken, N.J. and Kleist, J. (1993). The relationship of drought frequency and duration to time scales. AMS 8th Conference on Applied Climatology, (January), 179–184.
- Oliver, J.E. (1980). Monthly precipitation distribution: a comparative index. The Professional Geographer. 1980; 32(3):300–309.
- Sisay, E., Halefom, A., Khare, D., Singh, L. and Worku, T. (2017). Hydrological modelling of ungauged urban watershed using SWAT model. Modeling Earth Systems and Environment, 3(2), pp.693-702.

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Thank you for your attention!!