

Stochastic Lattice Models for Clouds and Parameterization of Organized Convection

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ICTP/PIMS summer school, July 1-19, 2024.

- ▶ Lecture 1: Stochastic modelling, Markov processes, Theory and Simulation (9 am - 10:30 am); Chapter 4 and Section 10.2.
- ▶ **Lecture 2: The stochastic multcloud model (11 am - 12:30 pm) ; Chapter 10: Sections 10.3 to 10.6**
- ▶ Hands-on activities (2:30 pm - 3:45 pm):
- ▶ Lecture 3: Waves and convective organization in the SMCM (4:00 pm - 5:15 pm) ; Chapter 11.

Introduction

The three-cloud type SMCM

Case when local interactions are ignored

Coupling the SMCM to a cumulus parameterization: Single Column Toy GCM

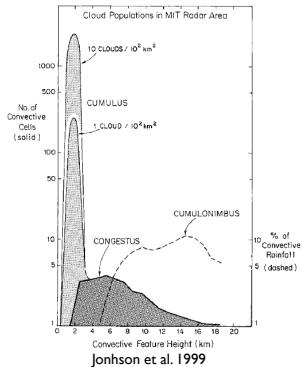
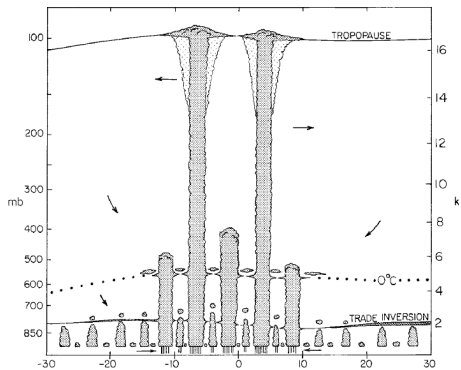
SMCM with local interactions

Parameter calibration using obs data a.k.a machine learning

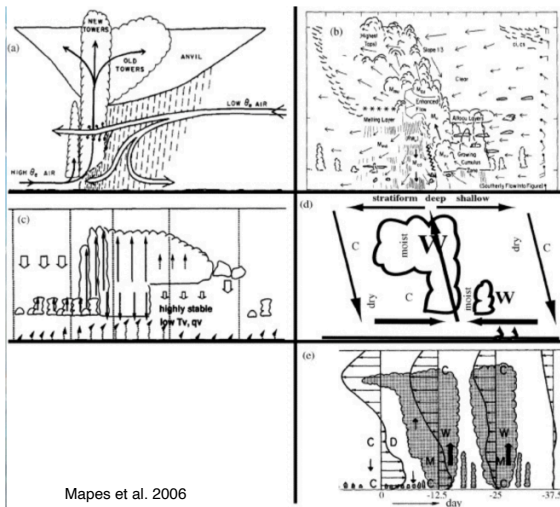
Introduction

- ▶ Organized convection in tropics involves three cloud types that interact with each other and with the environment
- ▶ Statistical-self similarity across-scales of tropical convective systems
- ▶ Deviation from quasi-equilibrium paradigm
- ▶ Impacts fidelity of GCMs with regards to tropical weather and climate
- ▶ Importance of missing subgrid variability due to organized convection as a multiscale process
- ▶ Design a stochastic models to mimic these subgrid scale interactions without substantial computational overhead in lieu of say super-parameterization

Tri-modal nature of tropical convection



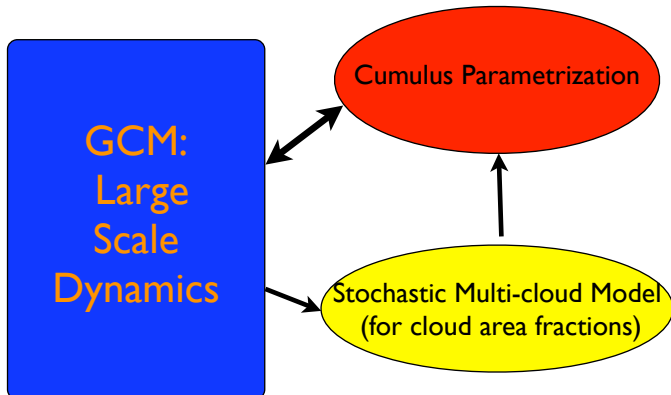
Self-similarity across scales



Mapes et al. 2006

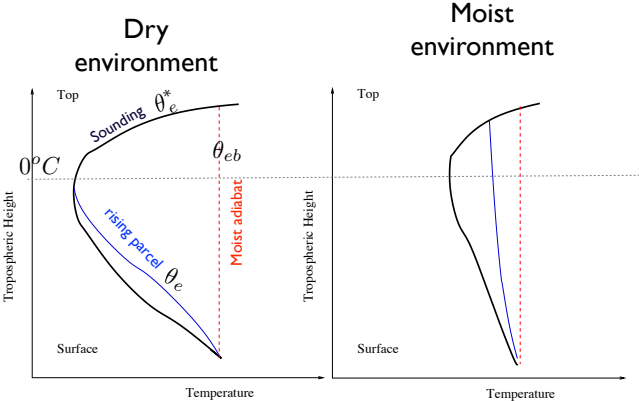
The stochastic model to add on an existing cumulus scheme

Stochastic Multi-cloud Model to inform cumulus parametrization: represent the missing sub-grid scale variability



Dilute Parcel lifting

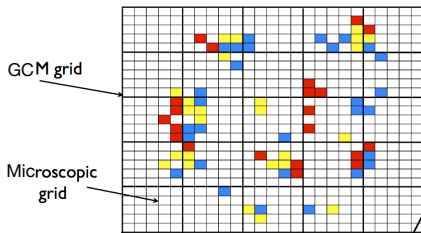
Dilute parcel lifting



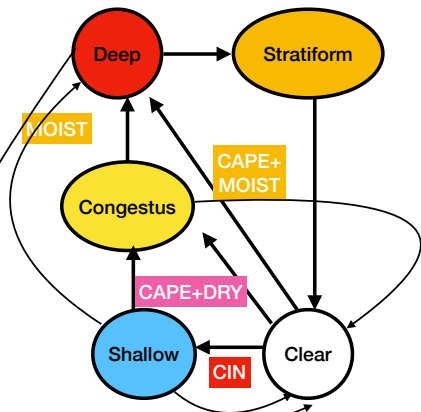
Multi-type Particle Microscopic Lattice Model

Multidimensional Markov process on a Lattice with Conditional Probability Rules

Fine lattice overlaid on GCM Grid



Each lattice is either occupied by a cloud of certain type or is clear sky... time varying!



The devil is in the details:

Play simple: The three cloud type SCM.

$$X_t^i = \begin{cases} 0 & \text{if site } i \text{ is clear sky} \\ 1 & \text{if site } i \text{ is occupied by a congestus cloud} \\ 2 & \text{if site } i \text{ is occupied by a deep convective cloud} \\ 3 & \text{if site } i \text{ is occupied by a stratiform anvil.} \end{cases} \quad (1)$$

$$P_{lk}^i \equiv \text{Prob} \{X_{t+\Delta t}^i = k / X_t^i = l\} = R_{lk}^i \Delta t + o(\Delta t), \quad (2)$$

for $l, k = 0, 1, 2, 3$, and $l \neq k$

$$P_{ll}^i \equiv \text{Prob} \{X_{t+\Delta t}^i = l / X_t^i = l\} = 1 - \sum_{k=0, k \neq l}^3 P_{lk}^i, \quad (3)$$
$$= 1 - \Delta t \sum_{k=0, k \neq l}^3 R_{lk}^i + O(\Delta t)$$

Intuitive Transition Rules

1. A clear site turns into a congestus site with high probability if CAPE is positive and the middle troposphere is dry.
2. A congestus or clear sky site turns into a deep convective site with high probability if CAPE is positive and the middle troposphere is moist.
3. A deep convective site turns into a stratiform site with high probability with a prescribed conversion rate, which may or may not depend on the state of the environment.
4. A cloudy site turns back to a clear sky with a certain probability according to a prescribed decay time scale for each cloud type.
5. It is very unlikely, during the short period of time Δt , for a clear sky or a congestus site to turn into a stratiform site, for a deep convective or stratiform site to turn into a congestus site, nor for a stratiform site to turn into a deep convective site.

The Matrix of transition rates

- ▶ Forbidden transitions

$$R_{03} = R_{13} = R_{21} = R_{31} = R_{32} = 0.$$

- ▶ Markov process at each site (independently on whether sites are connected or not)

$$R = \begin{bmatrix} -R_{01} - R_{02} & R_{01} & R_{02} & 0 \\ R_{10} & -R_{10} - R_{12} & R_{12} & 0 \\ R_{20} & 0 & R_{20} - R_{23} & R_{23} \\ R_{30} & 0 & 0 & -R_{30} \end{bmatrix}. \quad (4)$$

Case when local interactions are ignored

- ▶ Based on intuitive rules, we set

$$\begin{aligned}R_{01} &= \frac{1}{\tau_{01}} \Gamma(C) \Gamma(D), & R_{02} &= \frac{1}{\tau_{02}} \Gamma(C) (1 - \Gamma(D)), \\R_{10} &= \frac{1}{\tau_{10}} \Gamma(D), & R_{12} &= \frac{1}{\tau_{12}} \Gamma(C) (1 - \Gamma(D)), \\R_{20} &= \frac{1}{\tau_{20}} (1 - \Gamma(C)), & R_{23} &= 1/\tau_{23}, & R_{30} &= 1/\tau_{30}.\end{aligned}\tag{5}$$

where

$$\Gamma(x) \equiv \begin{cases} 1 - e^{-x} & \text{if } x > 0 \\ 0 & \text{Otherwise.} \end{cases}\tag{6}$$

- ▶ The transition time scales τ_{kl} are inferred from observations (radar data) using Bayesian machine learning: De La Chevrotiere et al. (2015,2016), Cardoso-Bihlo et al. (2019), ... **Carlos's talk next week!**

Academic example

Time	description	Case 1	Case 2
τ_{01}	formation of congestus	1 hour	3 hours
τ_{10}	decay of congestus	5 hours	2 hours
τ_{12}	conversion of congestus to deep	1 hour	2 hours
τ_{02}	formation of deep	2 hours	5 hours
τ_{23}	conversion of deep to stratiform	3 hours	0.5 hour
τ_{20}	decay of deep	5 hours	5 hours
τ_{30}	decay of stratiform	5 hours	24 hours

The stationary distribution, cloud area fractions, and the equilibrium statistics of the lattice model

► Equilibrium measure

$$\mathcal{P}_e R = 0 \quad (\text{Steady state of forward equations})$$

$$\mathcal{P}_e = \frac{1}{Z} \left(1, \frac{R_{01}}{R_{10}+R_{12}}, \frac{1}{R_{20}+R_{23}} \left(R_{02} + \frac{R_{12}R_{01}}{R_{10}+R_{12}} \right), \frac{R_{23}}{R_{30}} \frac{1}{R_{20}+R_{23}} \left(R_{02} + \frac{R_{12}R_{01}}{R_{10}+R_{12}} \right) \right) \quad (7)$$

► Cloud-type area fractions

$$\sigma_c = \frac{1}{N} \sum_{i=1}^N \mathbb{1}_{\{X_t^i=1\}}, \quad \sigma_d = \frac{1}{N} \sum_{i=1}^N \mathbb{1}_{\{X_t^i=2\}}, \quad \sigma_s = \frac{1}{N} \sum_{i=1}^N \mathbb{1}_{\{X_t^i=3\}} \quad (8)$$

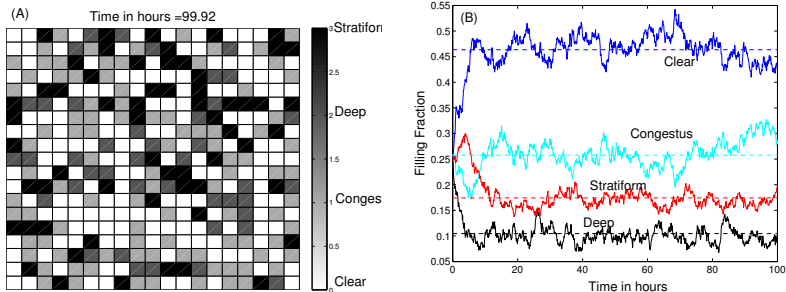
where

$$\mathbb{1}_{\{X_t^i=k\}} = \begin{cases} 1 & \text{if } X_t^i = k \\ 0 & \text{otherwise.} \end{cases}$$

The clear sky area fraction is given by

$$\sigma_{cs} \equiv 1 - \sigma_c - \sigma_d - \sigma_s.$$

Standalone example



Monte Carlo simulation SMCM with $n = 20$, $C = 0.25$, $D = 0.75$, Case 1. (A) A snapshot picture of one typical realization and (B) time series of the total coverages associated with each cloud type with the equilibrium values overlaid (dashed lines).

Coarse-grained process and mean-field equations

- ▶ Coarse grained process: $N = n \times n$ total number of lattice sites, N_c^t , N_d^t , and N_s^t are resp. the number of congestus, deep, and startiform sites at time t .
- ▶ Coverage fraction

$$\begin{aligned}\text{Prob}\{X_t^i = 1\} &= \frac{N_c^t}{N} = \sigma_c^t, & \text{Prob}\{X_t^i = 2\} &= \frac{N_d^t}{N} = \sigma_d^t, \\ \text{Prob}\{X_t^i = 3\} &= \frac{N_s^t}{N} = \sigma_s^t, & \text{Prob}\{X_t^i = 0\} &= \frac{N_{cs}^t}{N} = \sigma_{cs}^t, \quad \forall i = 1, 2, \dots, N.\end{aligned}\tag{9}$$

- ▶ Stochastic dynamics: Transition time is that of Min of i.i.d. Exp. R.V's:

$$\text{Prob}\{N_c^{t+\Delta t} = k + 1 / N_c^t = k\} = N_{cs}R_{01}\Delta t + o(\Delta t).\tag{10}$$

$$\text{Prob}\{N_c^{t+\Delta t} = k - 1 / N_c^t = k\} = N_c(R_{10} + R_{12})\Delta t + o(\Delta t),$$

$$\text{Prob}\{N_d^{t+\Delta t} = k + 1 / N_d^t = k\} = (N_{cs}R_{02} + N_cR_{12})\Delta t + o(\Delta t),$$

$$\text{Prob}\{N_d^{t+\Delta t} = k - 1 / N_d^t = k\} = N_d(R_{20} + R_{23})\Delta t + o(\Delta t),\tag{11}$$

$$\text{Prob}\{N_s^{t+\Delta t} = k + 1 / N_s^t = k\} = N_dR_{23}\Delta t + o(\Delta t),\tag{12}$$

$$\text{Prob}\{N_s^{t+\Delta t} = k - 1 / N_s^t = k\} = N_sR_{30}\Delta t + o(\Delta t),\tag{13}$$

- ▶ Mixed population birth and death process with immigration.

Probability transition matrix, Kolmogorov Equations, and Gillespie's algorithm

- ▶ 3D vector $Z_t = (N_c^t, N_d^t, N_s^t)$,
 $\epsilon_1 = (1, 0, 0)$, $\epsilon_2 = (0, 1, 0)$, $\epsilon_3 = (0, 0, 1)$.
- ▶ Kolmogorov backward Equations,

$$\begin{aligned} \frac{d}{dt} P_{i,j} = & N_{cs} R_{01} P_{i+\epsilon_1,j} + N_{cs} R_{02} P_{i+\epsilon_2,j} + N_c R_{12} P_{i+\epsilon_2-\epsilon_1,j} \\ & + N_d R_{23} P_{i+\epsilon_3-\epsilon_2,j} + N_c R_{10} P_{i-\epsilon_1,j} + N_d R_{20} P_{i-\epsilon_2,j} + N_s R_{30} P_{i-\epsilon_3,j} \\ & - [N_{cs}(R_{01} + R_{02}) + N_c(R_{10} + R_{12}) + N_d(R_{23} + R_{20}) + N_s R_{30}] P_{i,j} \\ & i, j \in \{0, 1, \dots, N\}^3. \end{aligned} \tag{14}$$

Solution: $P(t) = \exp\{Rt\} \in \mathcal{S} \subset \mathbb{R}^{N^3 \times N^3}$ (and full matrix)

- ▶ Efficiently evolved with Gillespie's algorithm (next slide)

Multidimensional Gillespie's exact algorithm:

- 0) Let $\lambda = N_{cs}(R_{01} + R_{02}) + N_c(R_{10} + R_{12}) + N_d(R_{23} + R_{20}) + N_s R_{30}$.
- 1) Let $Z_t = (N_c, N_d, N_s)$ be the state of the system at time t , $0 \leq t < \Delta T$.
- 2) Generate a random number r_1 uniformly from $(0, 1)$. Set $s = -\frac{1}{\lambda} \ln(r_1)$.
- 3) If $t + s > \Delta T$, then no transition occurs. Set $t = \Delta T$. Stop.
- 4) If $t + s \leq \Delta T$. Divide the interval into seven subintervals I_1, I_2, \dots, I_7 of sizes

$$\frac{N_{cs}R_{01}}{\lambda}, \frac{N_{cs}R_{02}}{\lambda}, \frac{N_cR_{10}}{\lambda}, \frac{N_cR_{12}}{\lambda}, \frac{N_dR_{23}}{\lambda}, \frac{N_dR_{20}}{\lambda}, \frac{N_sR_{30}}{\lambda},$$

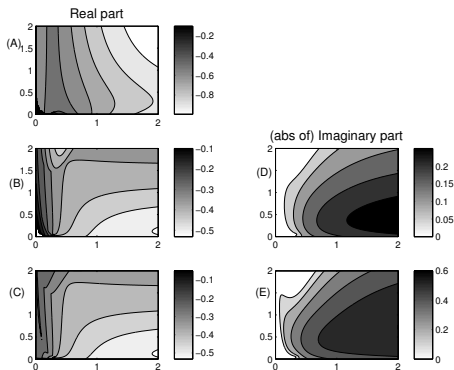
respectively.

- 5) Generate a random number r_2 . Select the subinterval I_k such that $r_2 \in I_k$. Then make the corresponding transition as follows.
 - ▶ If $r_2 \in I_1$, then $N_c = N_c + 1$.
 - ▶ If $r_2 \in I_2$, then $N_d = N_d + 1$.
 - ▶ If $r_2 \in I_3$, then $N_c = N_c - 1$.
 - ▶ If $r_2 \in I_4$, then $N_c = N_c - 1$, $N_d = N_d + 1$.
 - ▶ If $r_2 \in I_5$, then $N_d = N_d - 1$, $N_s = N_s + 1$.
 - ▶ If $r_2 \in I_6$, then $N_d = N_d - 1$.
 - ▶ If $r_2 \in I_7$, then $N_s = N_s - 1$.

- 6) Set $t = t + s$. If $t < \Delta T$ got to 1.

The mean field equations ... is damped

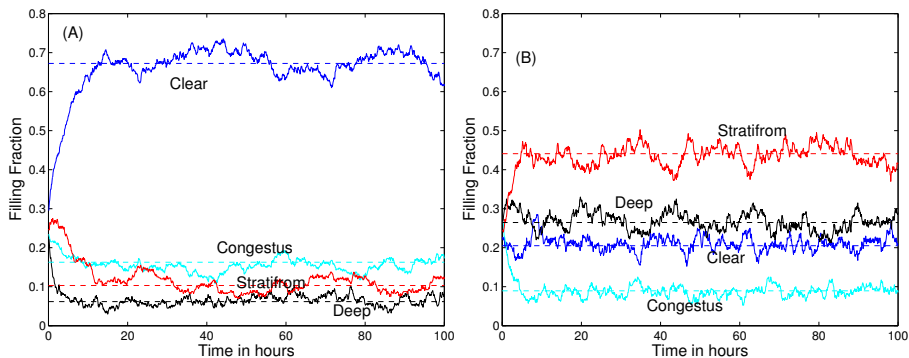
$$\begin{aligned}\dot{\sigma}_c &= (1 - \sigma_c - \sigma_d - \sigma_s)R_{01} - \sigma_c(R_{10} + R_{12}) \\ \dot{\sigma}_d &= (1 - \sigma_c - \sigma_d - \sigma_s)R_{02} + \sigma_c R_{12} - \sigma_d(R_{20} + R_{23}) \\ \dot{\sigma}_s &= \sigma_d R_{23} - \sigma_s R_{30}.\end{aligned}\quad (15)$$



Equilibrium eigenvalues of the mean field equations. (A),(B),(C): real part. (D): imaginary. (E):

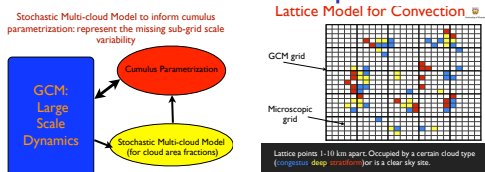
frequency/damping rate. CAPE C v.s dryness D on axes. **Without external forcing the system relaxes quickly to its equilibrium measure.**

Stochastic oscillations depend on frequency-to-damping ratio



Stochastic oscillations: (a) small (< 0.1) and (b) large (≈ 0.6) frequency to damping ratio

Coupling the SMCM to a cumulus parameterization



- ▶ SMCM simulates evolution of convective area fraction of key cloud types
- ▶ Coupled to mass flux scheme in a straightforward manner, with the goal to break the quasi-equilibrium paradigm:

$$M_U = \sigma_U W_U$$

(Peters et al. 2017; Dorrestijn et al. 2016)

- ▶ Here we will follow a simpler approach: different heating profiles associated with different cloud types, not necessarily well captured by existing mass-flux schemes.
- ▶ Success the Multi-cloud Model Paradigm in modelling Convectively Coupled Waves and MJO (both linear theory and non-linear simulations)

2 baroclinic mode Toy GCM

$$\mathbf{v} = \sqrt{2}\mathbf{v}_1 \cos(z) + \sqrt{2}\mathbf{v}_2 \cos(2z),$$

$$\theta = \sqrt{2}\theta_1 \sin(z) + 2\sqrt{2}\theta_2 \sin(2z).$$

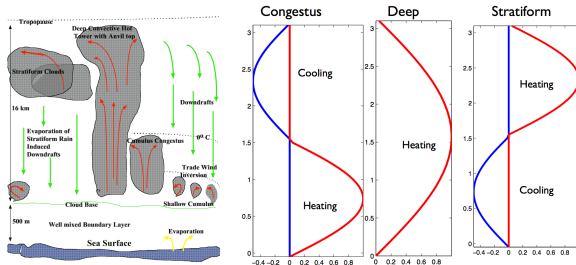
$$\frac{\partial \mathbf{v}_j}{\partial t} + \beta y \mathbf{v}_j^\perp - \nabla \theta_j = -C_d(u_0) \mathbf{v}_j - \frac{1}{\tau_W} \mathbf{v}_j, \quad j = 1, 2,$$

$$\frac{\partial \theta_1}{\partial t} - \text{div} \mathbf{v}_1 = \bar{H}_d + \xi_s \bar{H}_c + \xi_s \bar{H}_s + S_1$$

$$\frac{\partial \theta_2}{\partial t} - \frac{1}{4} \text{div} \mathbf{v}_2 = H_c - H_s + S_2,$$

$$\frac{\partial \theta_{eb}}{\partial t} = \frac{1}{h_b} (E - D)$$

$$\frac{\partial q}{\partial t} + \text{div} ((\mathbf{v}_1 + \tilde{\alpha} \mathbf{v}_2) \mathbf{q}) + \tilde{Q} \text{div} (\mathbf{v}_1 + \tilde{\lambda} \mathbf{v}_2) = -P + \frac{1}{H_T} D$$



Khouider and Majda, 2006; 2008: Simple models for convectively coupled waves

$$H_c = \sigma_c \frac{\bar{\alpha} \alpha_c}{H_m} \sqrt{CAPE_l^+}$$

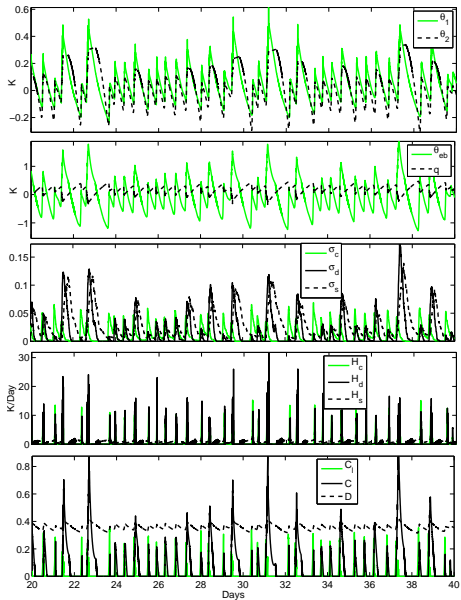
$$H_d = [\bar{\sigma}_d \bar{Q} + \frac{1}{\tau_c(\sigma_d)} (a_1 \theta_{eb} + a_2 q - a_0(\theta_1 + \gamma_2 \theta_2))]^+$$

$$\tau_c(\sigma_d) = \frac{\bar{\sigma}_d}{\sigma_d} \tau_c^0$$

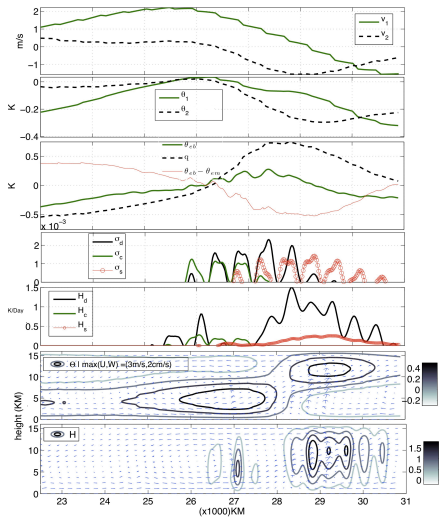
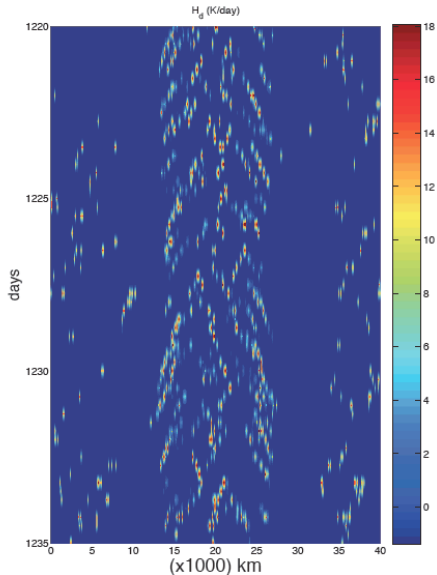
$$H_s = \alpha_s [\bar{\sigma}_s \bar{Q} + \frac{1}{\tau_c(\sigma_s)} (a_1 \theta_{eb} + a_2 q - a_0(\theta_1 + \gamma_2 \theta_2))]^+$$

$$\tau_c(\sigma_s) = \frac{\bar{\sigma}_s}{\sigma_s} \tau_c^0$$

Single column run



Waves along the equator with warm pool forcing



SMCM with local interactions

Why local interactions?

- ▶ Emulate self-organization of cloud clusters at the meso-scale: MCS's Estern Pacific ITCZ; Coastal convection, etc.
- ▶ Represent unresolved effects of shear, sea-breeze, orography, etc. on organization of convection, turbulent mixing, radiation feedback, etc.
- ▶ Cloud-cloud interactions: can be both favourable and unfavourable to organization

The multi-type particle interacting system

- ▶ Multiple particle Hamiltonian

$$H(X, U) = \sum_{k=1}^3 \sum_{l=k}^3 E_{k,l}(X) + g_k(U)X$$

$$E_{k,l}(X) = - \sum_{i,j=1, i \neq j}^N J_{k,l}(|i-j|) \mathbb{1}_{\{X^i=k\}} \mathbb{1}_{\{X^j=l\}}, 1 \leq k \leq l \leq 3,$$

$$J_{k,l}(r) \neq 0 \iff r = 1 \text{ or } 0 < r \leq \sqrt{2},$$

- ▶ Choose Transition Rates to yield Gibbs equilibrium measure:

$$\begin{aligned} R_{02} + R_{01} &= \frac{\rho_1}{\rho_0} R_{10} e^{H_0 - H_1} + \frac{\rho_2}{\rho_0} R_{20} e^{H_0 - H_2} + \frac{\rho_3}{\rho_0} R_{30} e^{H_0 - H_3} \\ R_{01} &= \frac{\rho_1}{\rho_0} (R_{12} + R_{10}) e^{H_0 - H_1} \\ R_{02} &= \frac{\rho_2}{\rho_0} (R_{20} + R_{23}) e^{H_0 - H_2} - \frac{\rho_1}{\rho_0} R_{12} e^{H_0 - H_1} \\ R_{30} &= \frac{\rho_3}{\rho_2} R_{23} e^{H_2 - H_3}, \forall j = 1, \dots, N, \forall X \in \Sigma. \end{aligned} \tag{16}$$

ρ is the "background" distribution counting for external potential, associated with background rates $\tilde{R}_{kl}(U)$. $H_k - H_j$ are local-interaction energy difference. [Partial detailed balance.](#)

- ▶ Equilibrium measure:

$$\mu(X_t) \propto e^{-H(X_t)}.$$

Coarse grained approximation

- ▶ Coarse-grained Hamiltonian as a conditional expectation:

$$\bar{H}(\bar{X}) = E[H(X)|\bar{X}] = \sum_{k=1}^3 \sum_{l=1}^3 E[E_{kl}(X)/\bar{X}] + \sum_{k=1}^3 g_k(U) \sum_{i=1}^M N_k^i,$$

$$E[E_{kl}(X)/\bar{X}] = - \sum_{k=1}^3 \sum_{j=1}^N \sum_{l=1}^3 \sum_{r=1, r \neq j}^N E \left[J_{kl}(|j-r|) \mathbb{1}_{\{X_j=k\}} \mathbb{1}_{\{X_j=l\}} / \bar{X} \right].$$

- ▶ Uniform redistribution of particles, yields

$$E[E_{kl}(X)/\bar{X}] = - \frac{1}{Q^2} \sum_{k=1}^3 \sum_{l=1}^3 \sum_{i=1}^M J_{kl}^0 \left[(n_b(q-2)^2 + 4(n_b - n'_b)(q-2) + 4n_1) N_i^j \right. \\ \left. + (n'_b(q-2) + n_2) (N_i^{js} + N_i^{jn} + N_i^{je} + N_i^{jw}) + n_3 (N_i^{jsw} + N_i^{jse} + N_i^{jnw} + N_i^{jne}) \right] N_k^i,$$

NW	N	NE
W		E
SW	S	SE

Retains interactions across coarse cells.

Mean field equations

In the limit of a large micro and macro lattice sizes we get

$$\begin{aligned}\frac{\partial \sigma_1}{\partial t} &= \alpha_{01} \sigma_0 - (\alpha_{10} + \alpha_{12}) \sigma_1 \\ \frac{\partial \sigma_2}{\partial t} &= \alpha_{02} \sigma_0 + \alpha_{12} \sigma_1 - (\alpha_{20} + \alpha_{23}) \sigma_2 \\ \frac{\partial \sigma_3}{\partial t} &= \alpha_{23} \sigma_2 - \alpha_{30} \sigma_3,\end{aligned}\tag{17}$$

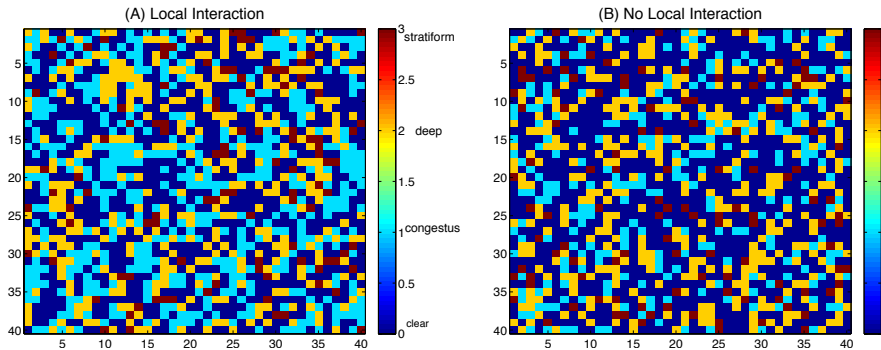
with

$$\begin{aligned}\alpha_{k0} &= \tilde{R}_{k0}, \quad k = 1, 2, 3, \quad \alpha_{01} = \tilde{R}_{01} e^{\Gamma_{10}}, \quad \alpha_{12} = \tilde{R}_{12} e^{\Gamma_{12}}, \quad \alpha_{23} = \tilde{R}_{23} e^{\Gamma_{23}}, \\ \alpha_{02} &= \frac{1}{\rho_0} (\rho_2 \tilde{R}_{20} - \rho_1 \tilde{R}_{12}) e^{\Gamma_{20}} + \frac{\rho_3}{\rho_0} \tilde{R}_{30} e^{\Gamma_{30}}\end{aligned}\tag{18}$$

$$\Gamma_{0k}(x) = \sum_{l=1}^3 J_{kl} * \sigma_l(x) \text{ and } \Gamma_{kl}(x) = \Gamma_{0l}(x) - \Gamma_{0k}(x),\tag{19}$$

with $f * g(x) = \int f(x-y)g(y) dy$ is the convolution operation.

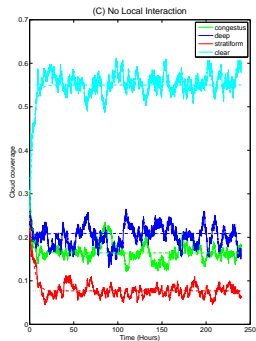
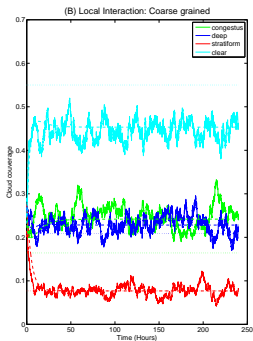
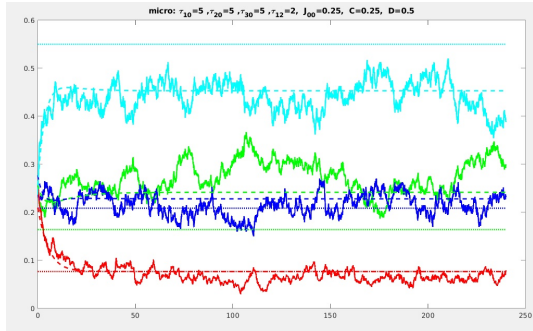
Numerical tests



Snapshot of the multicloud microscopic lattice model. (a) With local interactions and (b) without local interactions. $n = 40$.

$$\text{Interaction matrix } J_0 = \begin{bmatrix} 0.25 & 0 & 0 \\ 0 & 0.125 & 0.05 \\ 0 & 0.05 & 0.125 \end{bmatrix}$$

dash: Mean field
dots: background

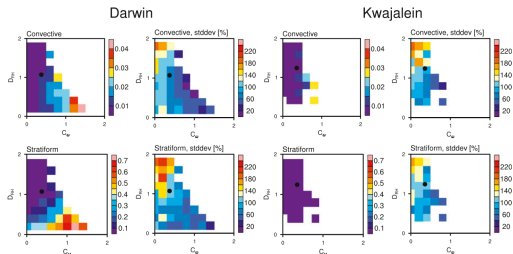


Parameter calibration using obs data

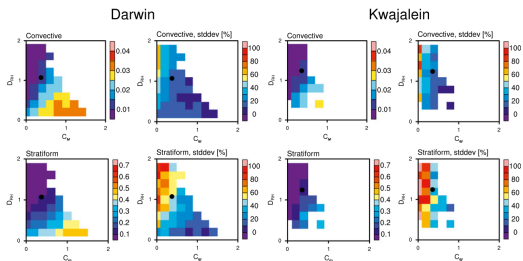
- ▶ Transition probabilities of cloud types are uncertain
- ▶ Use brute-force computations using LES data. (Dorrestijn et al. 2016)
- ▶ The formalism allows to reduce the task to learning key parameters: transition times and CAPE0, CIN0, reference values, etc.
- ▶ What is dequate cost function:
 - ▶ Cheap and dirty: Constraint the equilibrium distribution (Peters et al. 2017)
 - ▶ Precise but highly expensive
 - ▶ Midway pathways exists: Carlos Sevilla's thesis

Data from Darwin and Kwajalein

Radar data



SMCM



Regimes of forced v.s random convection: regions of dry and/low instability tend to display high CAF variance and low mean CAF

The Bayesian Inference Procedure

De La Chevrotiere et al. (2016; 2017)

- ▶ Find a distribution of the parameters

$$\boldsymbol{\theta} = (\tau_{01}, \tau_{10}, \tau_{12}, \tau_{02}, \tau_{23}, \tau_{20}, \tau_{30}).$$

given convective area fraction data and large scale predictors
— time series:

$$\mathbf{x}_t, \mathbf{u}_t, 1 \leq t \leq T$$

Prior distribution: $\pi(\boldsymbol{\theta}) \longrightarrow$ posterior distribution: $\pi(\boldsymbol{\theta}|\mathbf{x}_t)$

Data information encoded model *likelihood*: $f(\mathbf{x}_t|\boldsymbol{\theta}; \mathbf{u}_t)$;

- ▶ Bayes's theorem:

$$\pi(\boldsymbol{\theta}|\mathbf{x}_t) = \frac{f(\mathbf{x}_t|\boldsymbol{\theta})\pi(\boldsymbol{\theta})}{\int f(\mathbf{x}_t|\boldsymbol{\theta})\pi(\boldsymbol{\theta})d\boldsymbol{\theta}}. \quad (20)$$

Computing the likelihood

- ▶ Markov process

$$\begin{aligned} f(\mathbf{x}_{1:T} | \mathbf{u}_{1:T}, \boldsymbol{\theta}) &= \prod_{t=1}^T f_{t-1}(\mathbf{x}_t | \mathbf{x}_{t-1}, \mathbf{u}_{t-1}, \boldsymbol{\theta}) \\ &= \prod_{t=1}^T \mathbb{1}_{\phi(\mathbf{N}_c^{t-1}, \mathbf{N}_d^{t-1}, \mathbf{N}_s^{t-1})}^* \exp[R(\mathbf{u}_{t-1}, \boldsymbol{\theta})\Delta t] \mathbb{1}_{\phi(\mathbf{N}_c^t, \mathbf{N}_d^t, \mathbf{N}_s^t)} \end{aligned} \quad (21)$$

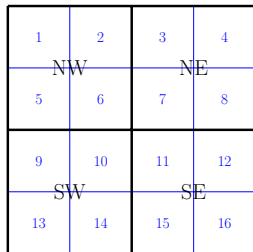
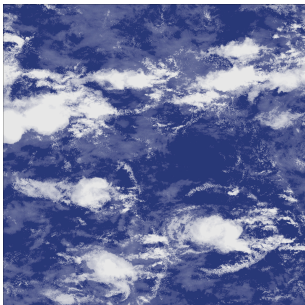
ϕ is the injection of the state space $\mathcal{S} \subset \mathbb{R}^3$ into the integers.

- ▶ MCMC allows sampling of posterior without knowing it explicitly.
- ▶ Computational burden is in exponential matrix:
 $\exp[R(\mathbf{u}_{t-1}, \boldsymbol{\theta})\Delta t]$
- ▶ Uniformization technique, stable and accurate but still inefficient!
- ▶ Convergence of MCMC can also be slow
- ▶ Alternatives: Variational inference in lieu of MCMC;
- ▶ There are also way to approximate f ... sample statistics

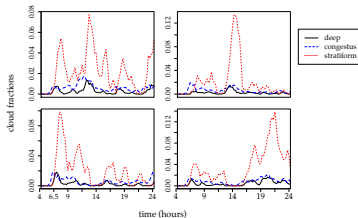
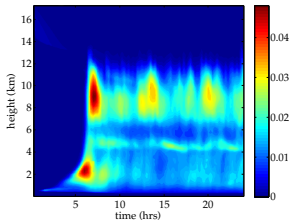
Giga LES data

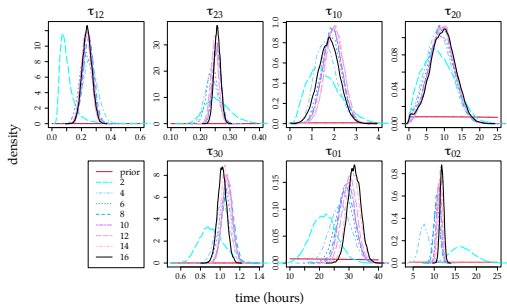
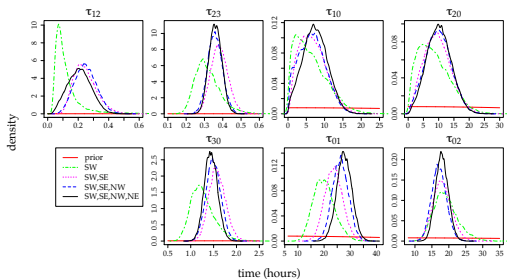
Khairoutdinov et al. (2001).

24 hr simulation $2048 \times 2048 \times 256$ grid; Domain size: $204.8 \times 204.8 \text{ km}^2 \times 27 \text{ km}$; data saved every 15 minutes.



cloud liquid/ice water mixing ratio ($\times 10^{-3} \text{ kg kg}^{-1}$)





Sequence of posterior distributions based on sequential Bayesian inference using a 2×2 (top set of panels) and 4×4 domain partitions (bottom panels)

Comparison of inferred time scales

Parameter	Mean (SD) [hours]		
	2 × 2 Partition	4 × 4 Partition	Peters et al.
τ_{01}	27.686 (8.233)	31.789 (4.795)	1
τ_{10}	7.426 (11.155)	1.761 (0.224)	1
τ_{12}	0.208 (0.006)	0.238 (0.001)	3
τ_{02}	17.950 (3.507)	11.821 (0.211)	4
τ_{23}	0.359 (0.001)	0.2570 (0.0001)	0.13
τ_{20}	10.126 (15.674)	9.551 (13.146)	5
τ_{30}	1.444 (0.021)	1.021 (0.002)	5