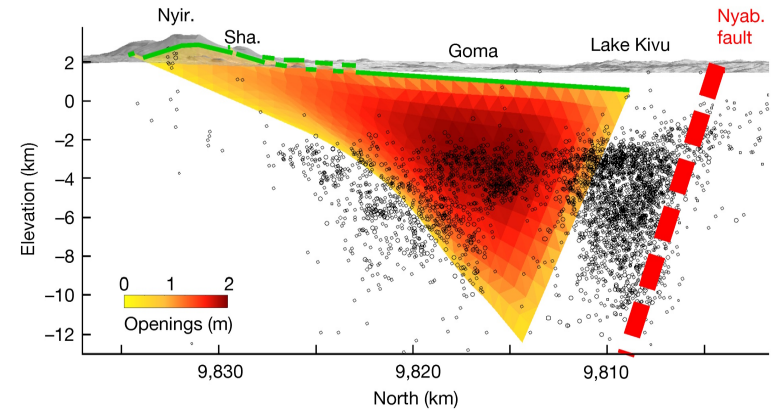
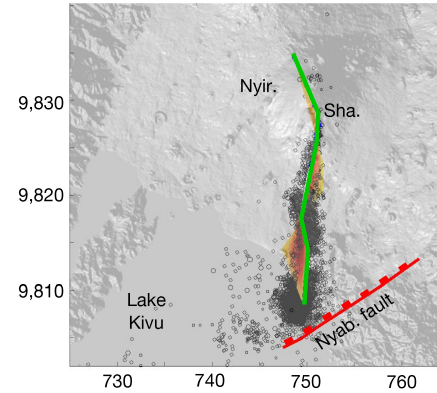
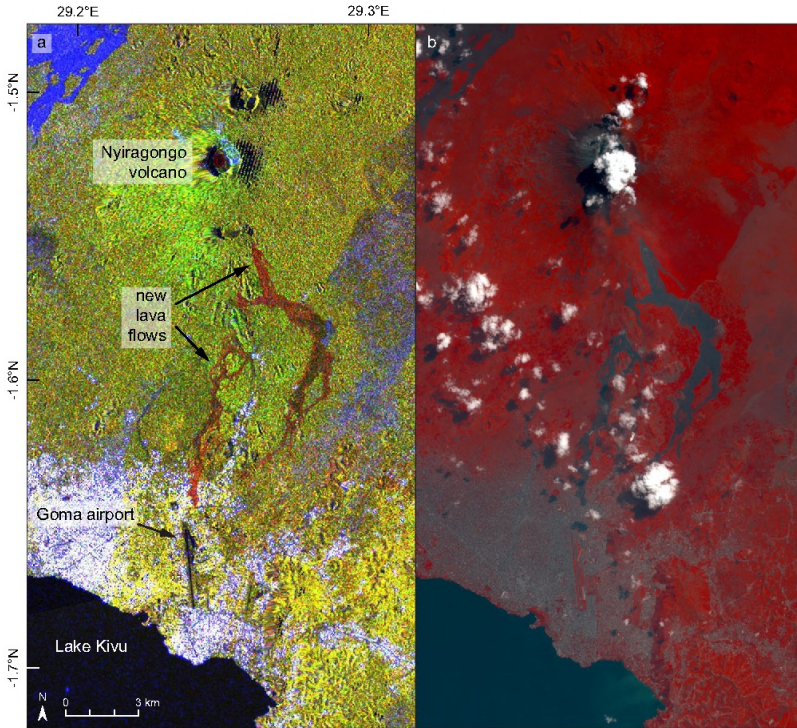


Towards a better understanding of subsurface magma propagation using numerical and analogue models, application to recent eruptions of Nyiragongo volcano

Virginie Pinel, C. Mériaux, S. Furst, F. Maccaferri, A. Gaete
& M. M. H. Mohamed

Nyiragongo volcano, DRC: 20 km of lateral propagation

1977, Jan. 2002, May 2021

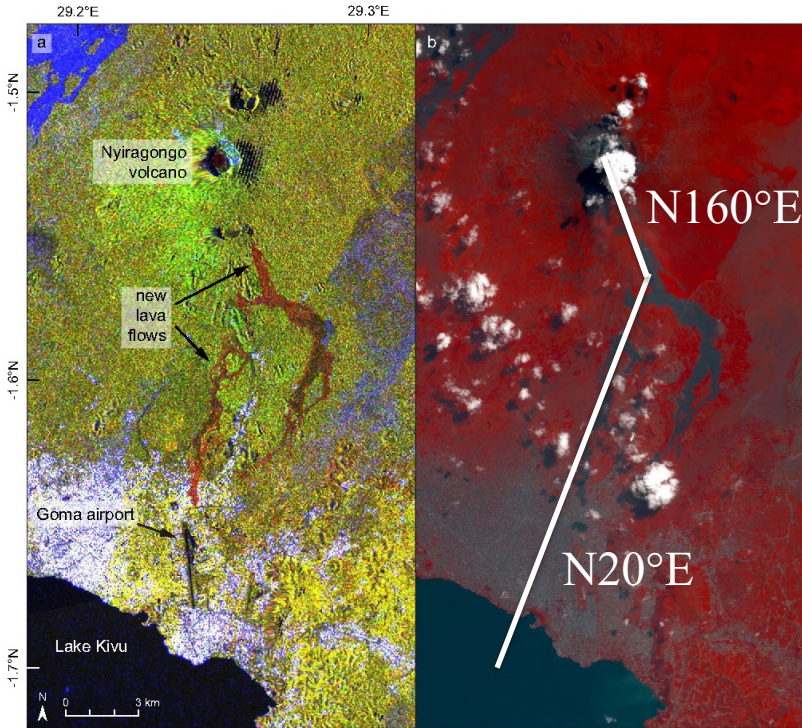


From Smittarello et al, Nature, 2022

From Poland, 2022

Nyiragongo volcano, DRC: 20 km of lateral propagation

1977, Jan. 2002, May 2021



From Poland, 2022

Key questions:

will the magma reach the surface?

where?

when?

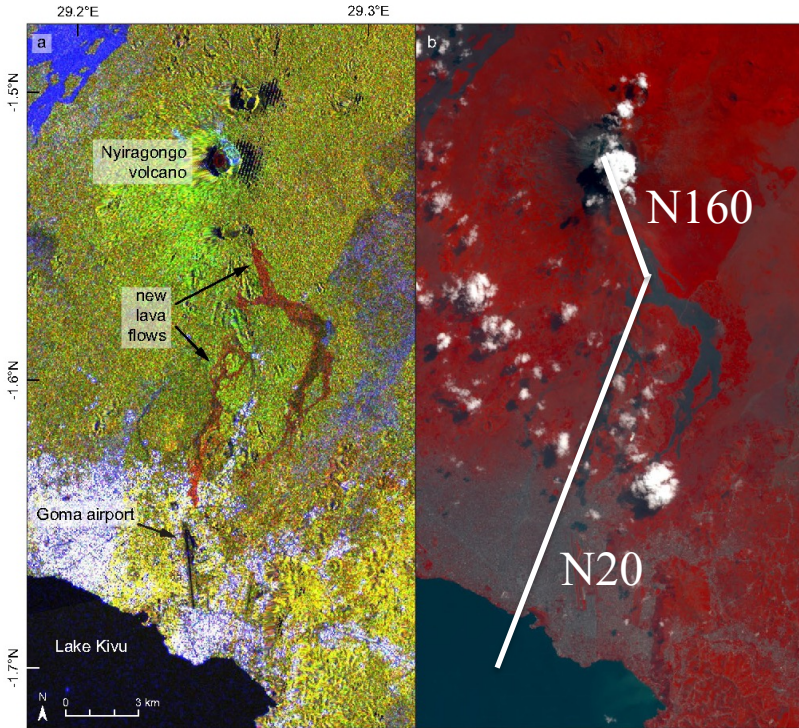
1. What controls the magma path and velocity?

2. Direction of propagation

→ lateral versus vertical

Nyiragongo volcano, DRC: 20 km of lateral propagation

1977, Jan. 2002, May 2021



From Poland, 2022

Key questions:

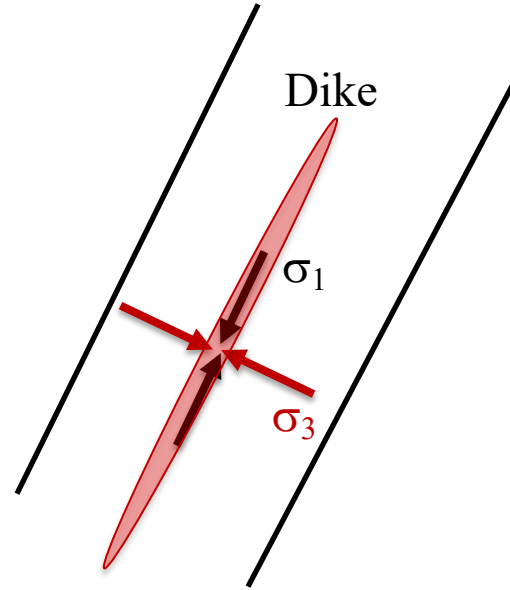
**will the magma reach the surface?
where?
when?**

**1. What controls the magma path
and velocity?**

**2. Direction of propagation
→ lateral versus vertical**

Dikes orientation is function of the local stress field

When the deviatoric stress field is large dykes should open perpendicular to the minimum compressive stress σ_3 (Anderson, 1951)



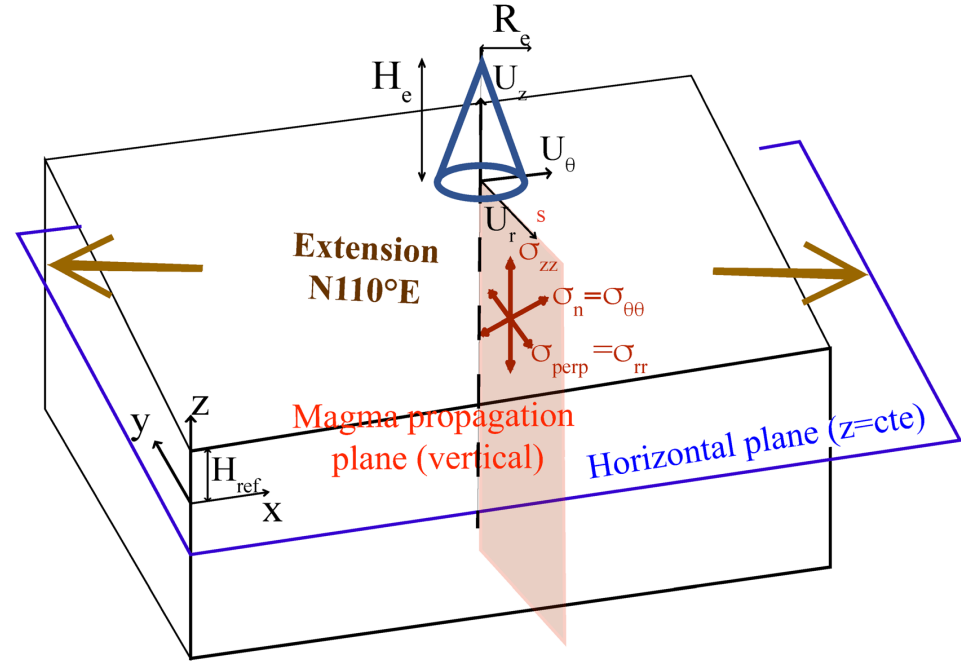
Modeling strategy for Nyiragongo case

3D stress field calculation:

- Lithostatic reference state
- Extension (N110°E)
- Edifice load

*Stress in horizontal plane
--> Dyke orientation

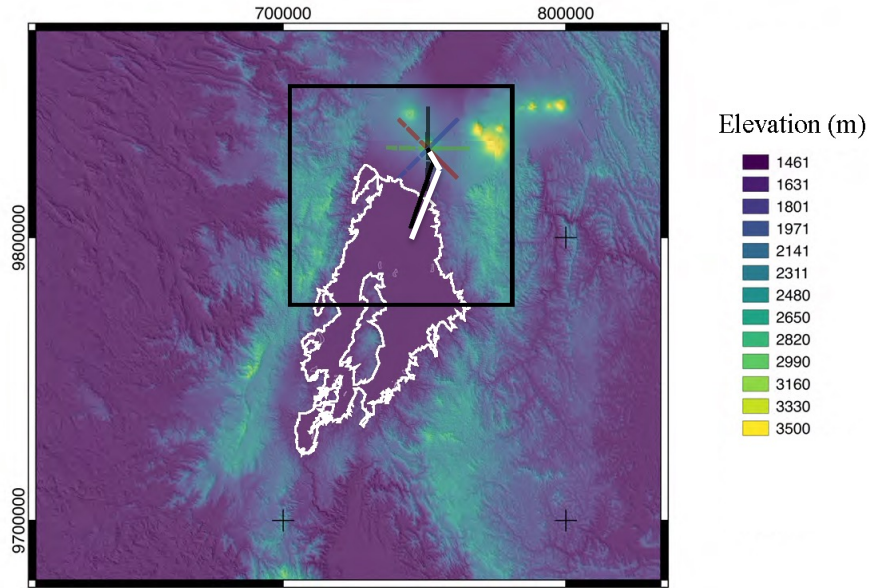
*Stress in vertical plane
--> Direction of propagation



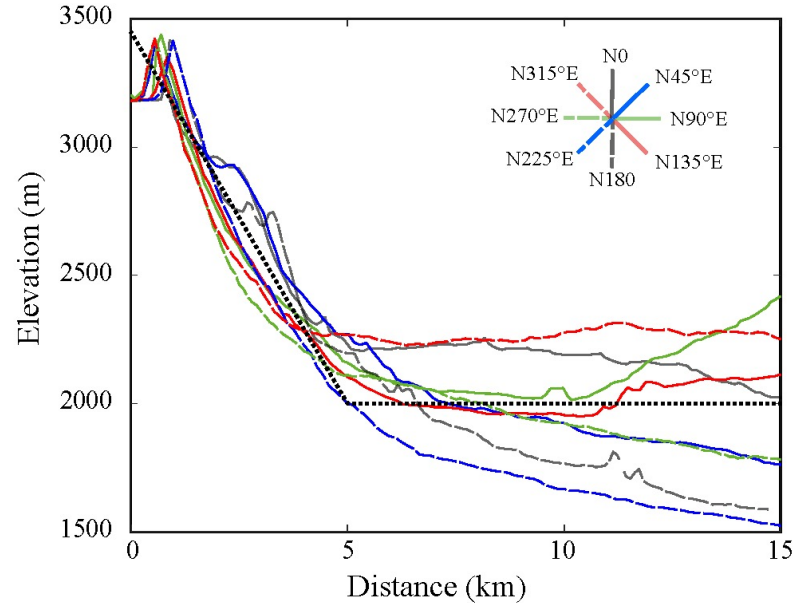
Analytical solution

Conical load applied at $H_{ref}=2\text{km}$

Nyiragongo volcano topography



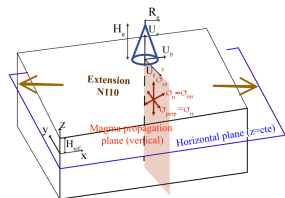
Nyiragongo edifice: 3470 m a.s.l.
Maximum water depth in Lake Kivu: 475 m



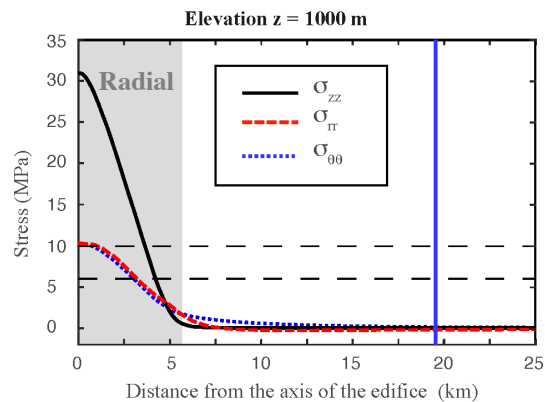
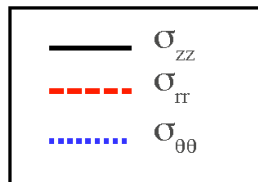
Nyiragongo topography
can be approximated by a **conical load**
(radius $R_e=5$ km, $H_e=1450$ m lying above $H_{ref}=2000$ m)

Topography decreases more towards the South, South-West

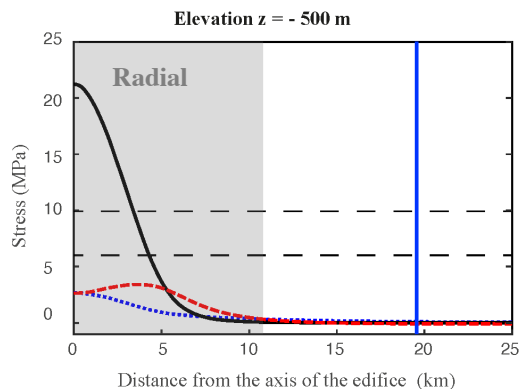
Edifice load effect is dominant at shallow level beneath the volcano



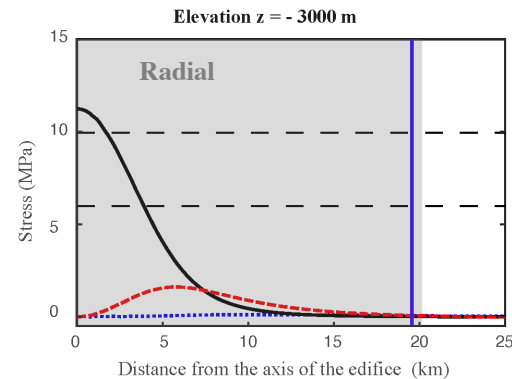
Analytical solution



Depth = $0.2 \cdot R_e$



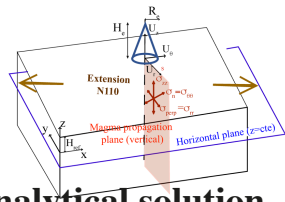
Depth = $0.5 \cdot R_e$



Depth = R_e

Increasing depth

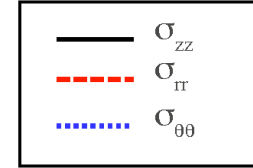
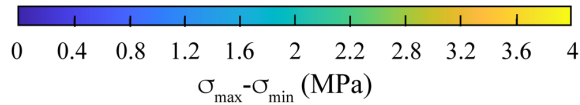
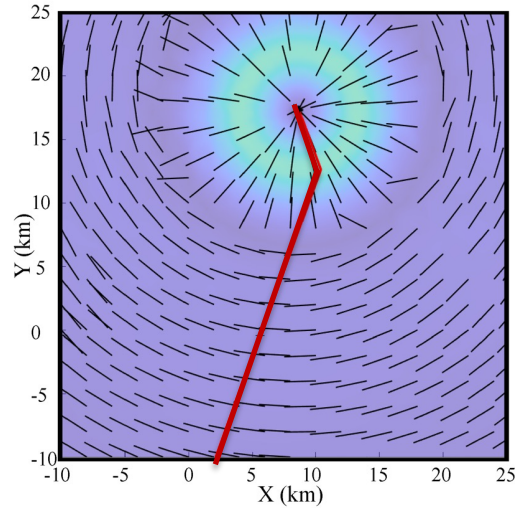
Dyke orientation: Maximum compressive stress direction in horizontal planes



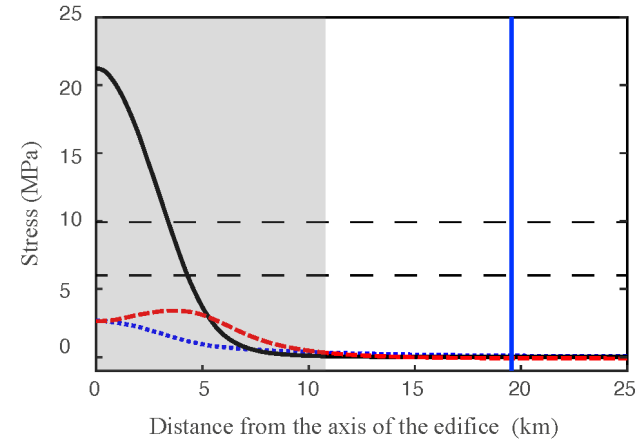
Analytical solution

No extension

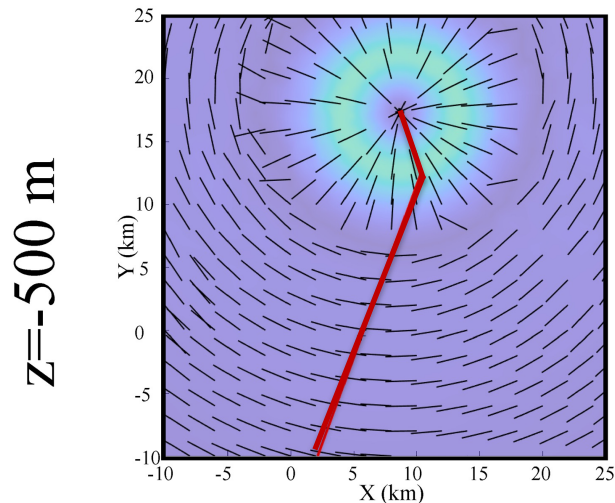
$z = -500 \text{ m}$



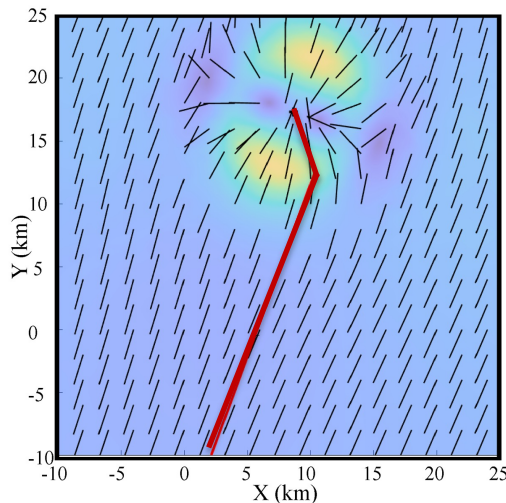
Elevation $z = -500 \text{ m}$



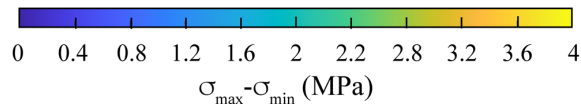
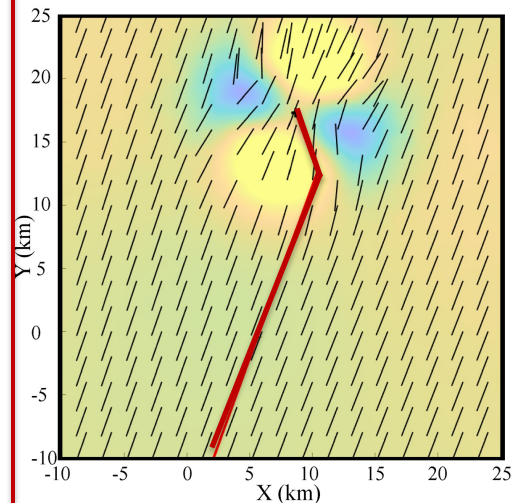
No extension


$$z = -500 \text{ m}$$

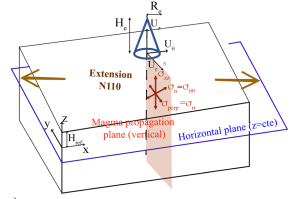
Uniform Extension N110°E, $\sigma=1$ MPa



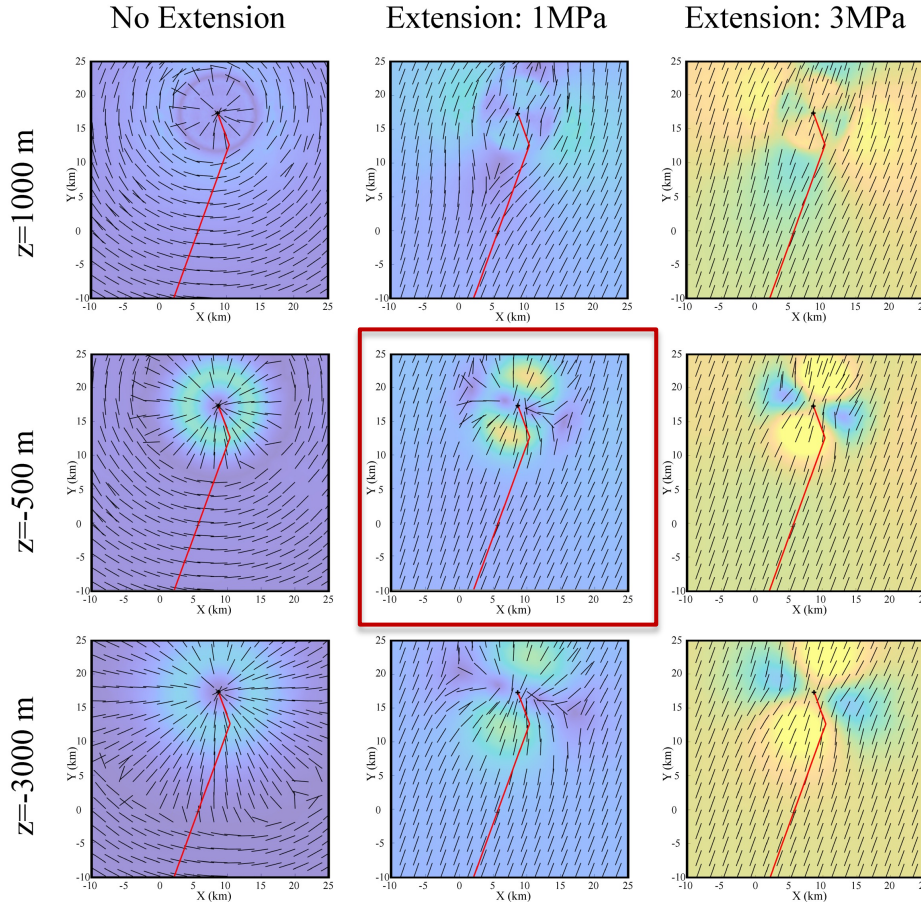
Uniform Extension N110°E, $\sigma=3\text{MPa}$



Dyke orientation: Maximum compressive stress direction in horizontal planes



Analytical solution

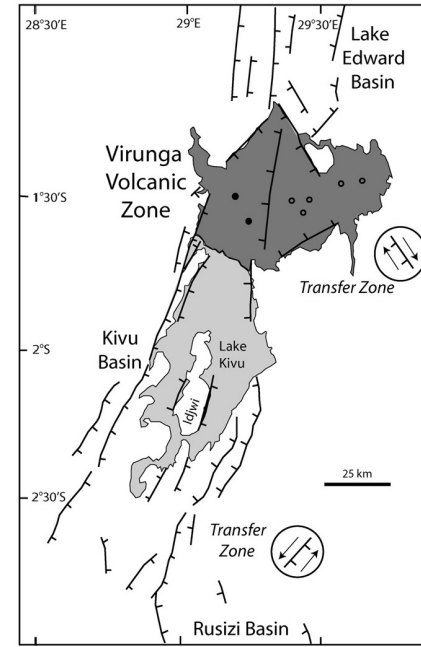
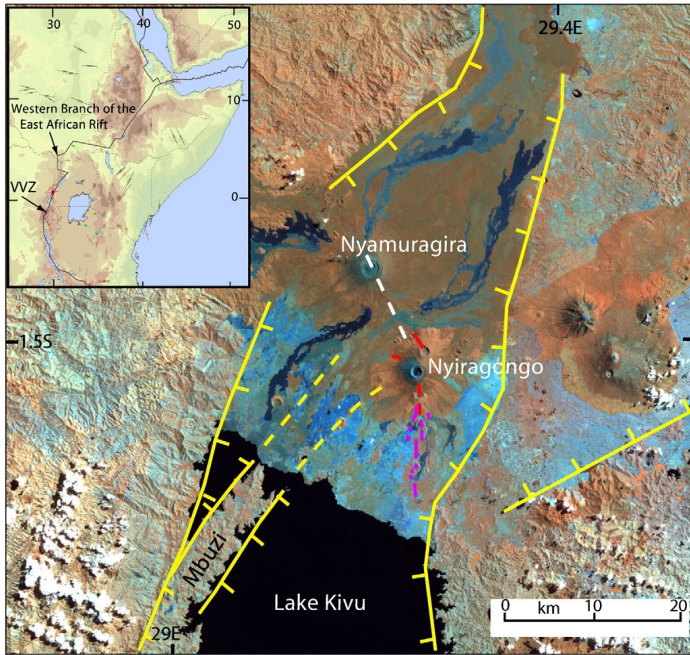


An extension of 0.5 to 2 MPa,
at a depth of between
500 m and -3000 m

→ Consistent dyke orientation

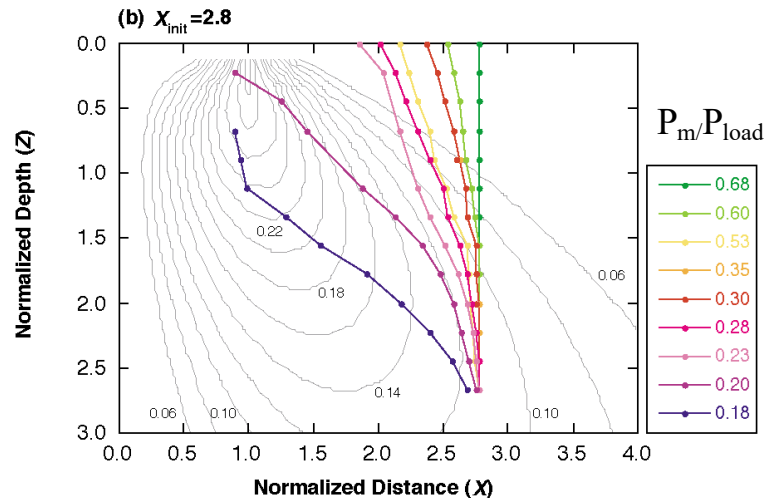
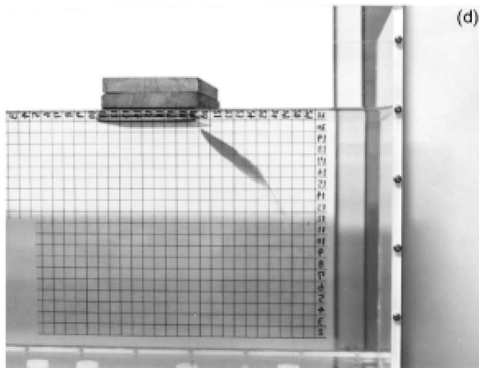
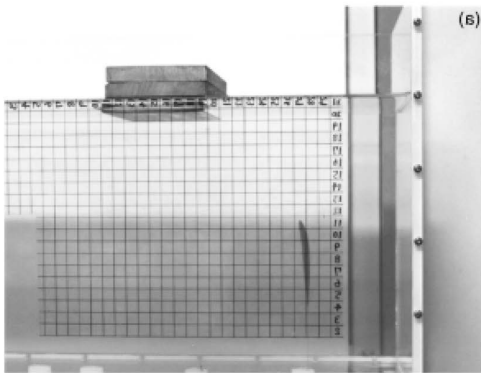
But tectonic stress is more complex

Local faults



From Wadge & Burt, 2011

Orientation also depends on the magma driving pressure



Local stress field influences the magma path: $R = \sigma_{xy}/P_m$

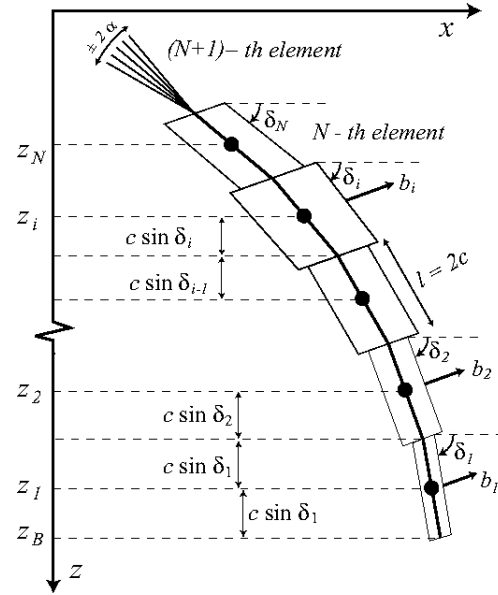
If $R > 0.2$ deflection

2D quasi-static model $\rightarrow$ path

A buoyant isolated crack modelled by BEM

The orientation that maximizes
the total energy release is chosen.

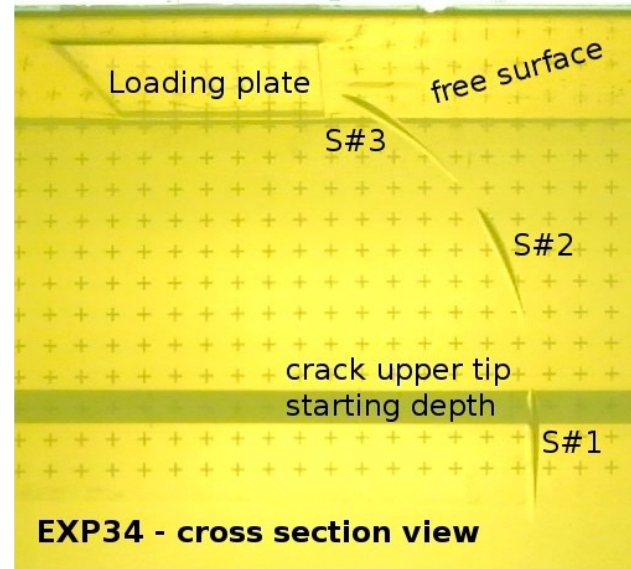
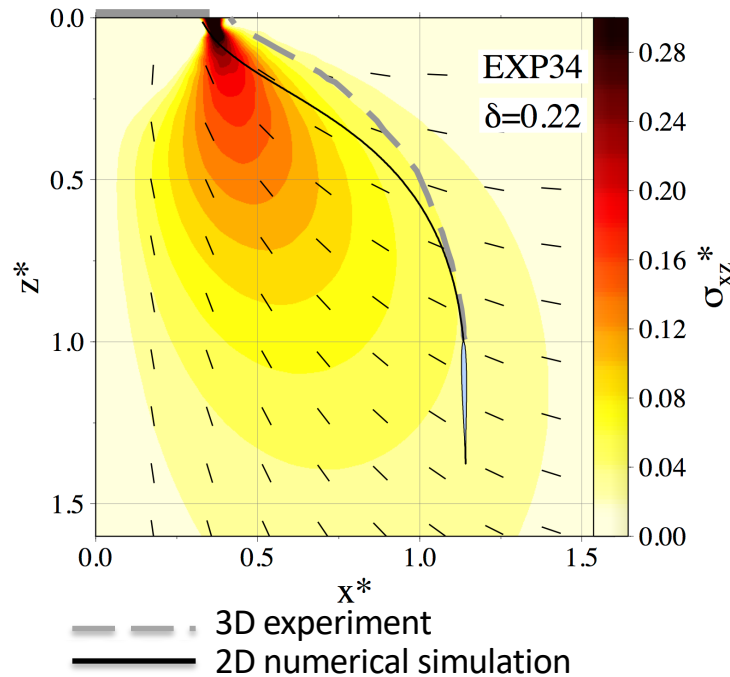
(strain + gravitational energy)



From Maccaferri et al, JVGR, 2011

Comparing numerical and analogue models

Criterion $\delta = \sigma_{xy} / (P_m * L^*)$

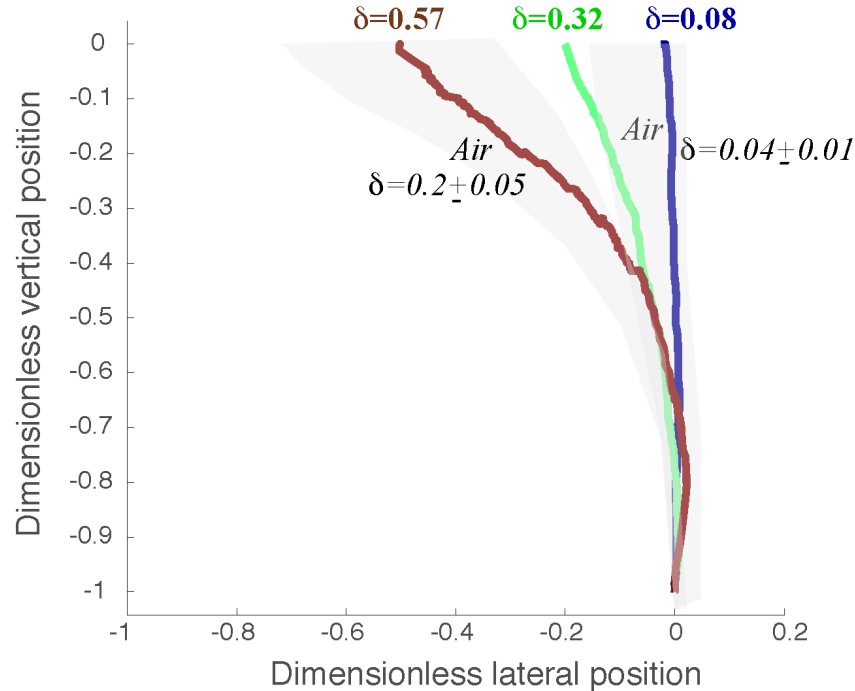


From Maccaferri et al, 2019

Influence of fluid viscosity...

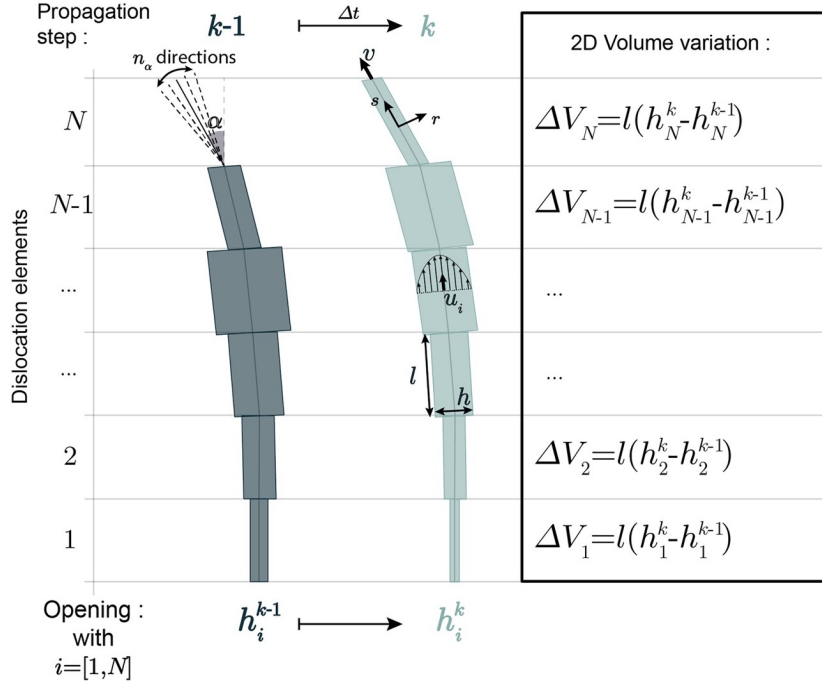
Injection of silicon oil
50 000 more viscous than air

Criterion $\delta = \sigma_{xy} / (P_m * L^*)$



Deflection decreases when viscosity increases

New dynamical models with full coupling

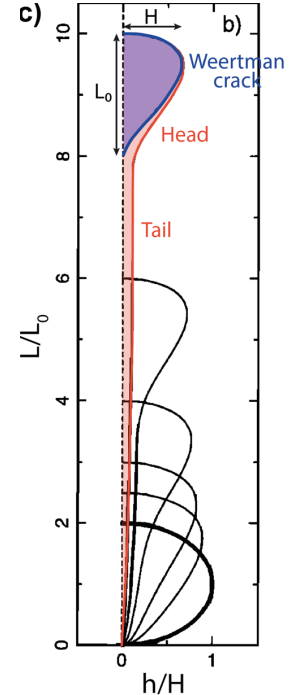


$$\Delta E = \Delta E_f \cdot l + \Delta E_v$$

$$v = \frac{\Delta E - \Delta E_f \cdot l}{\eta \cdot D'}$$

With

$$\Delta E_f = K_c^2 \frac{1 - nu}{2G}$$

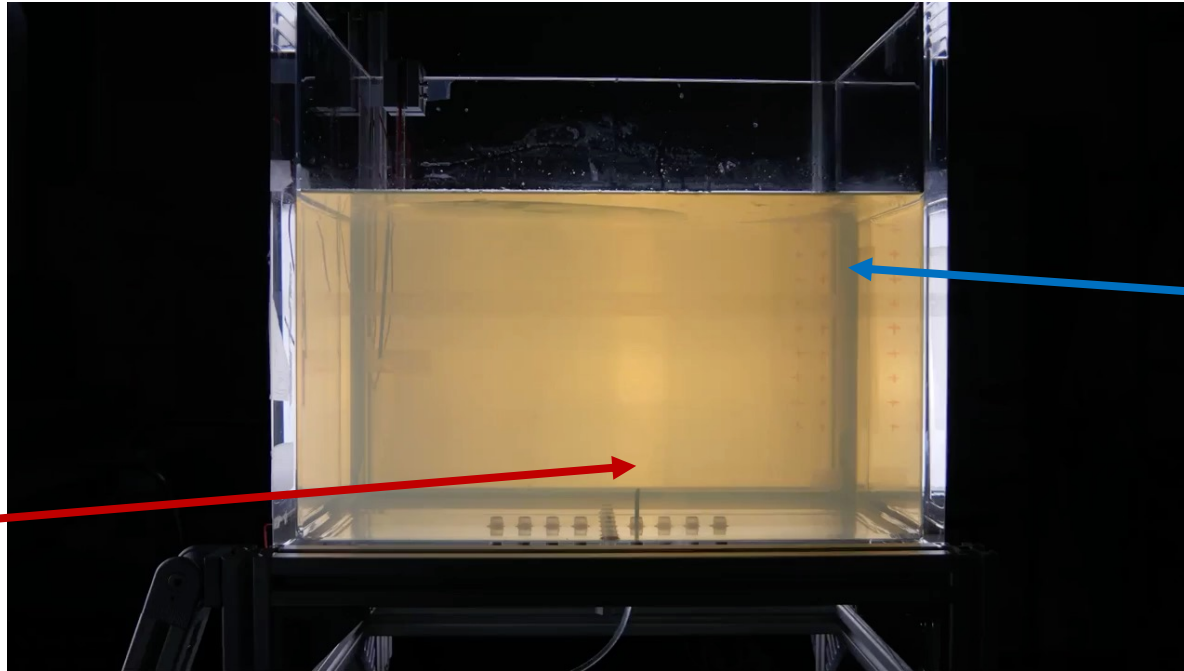


Validated with analytical solutions
for the viscous contribution

(strain + gravitational energy + **viscous energy**
→ **velocity**)

Validation by oil injection inside gelatin

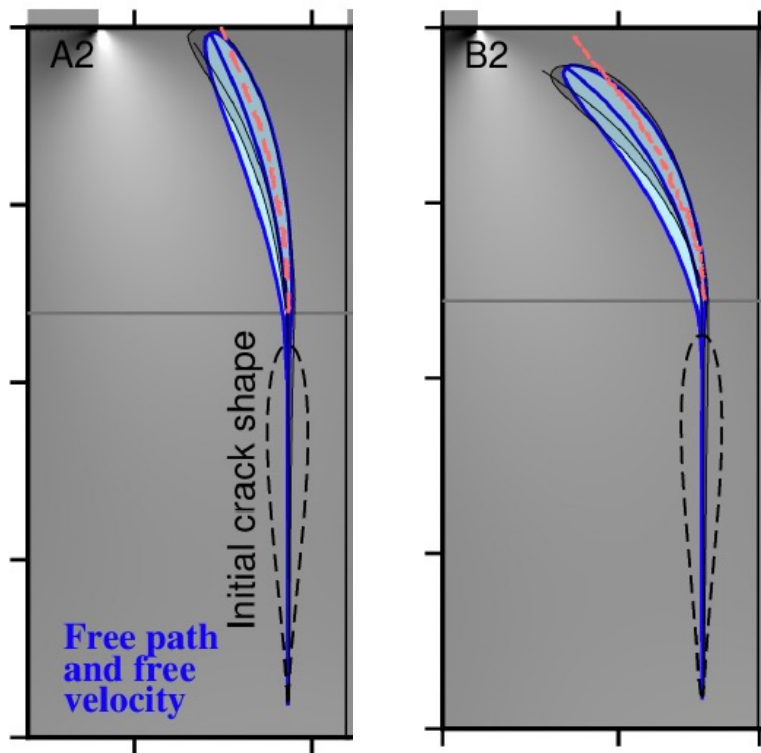
→ trajectories, velocities



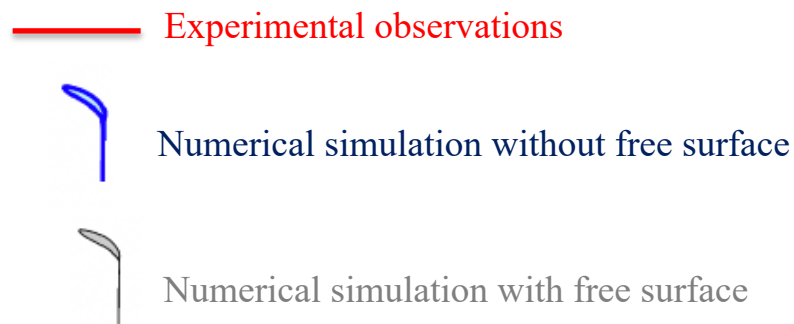
Fluid
(μ , ρ)
Silicon oil/magma

Elastic medium
(E , ν , K_c , ρ)
Salty gelatin/crust

Path consistency

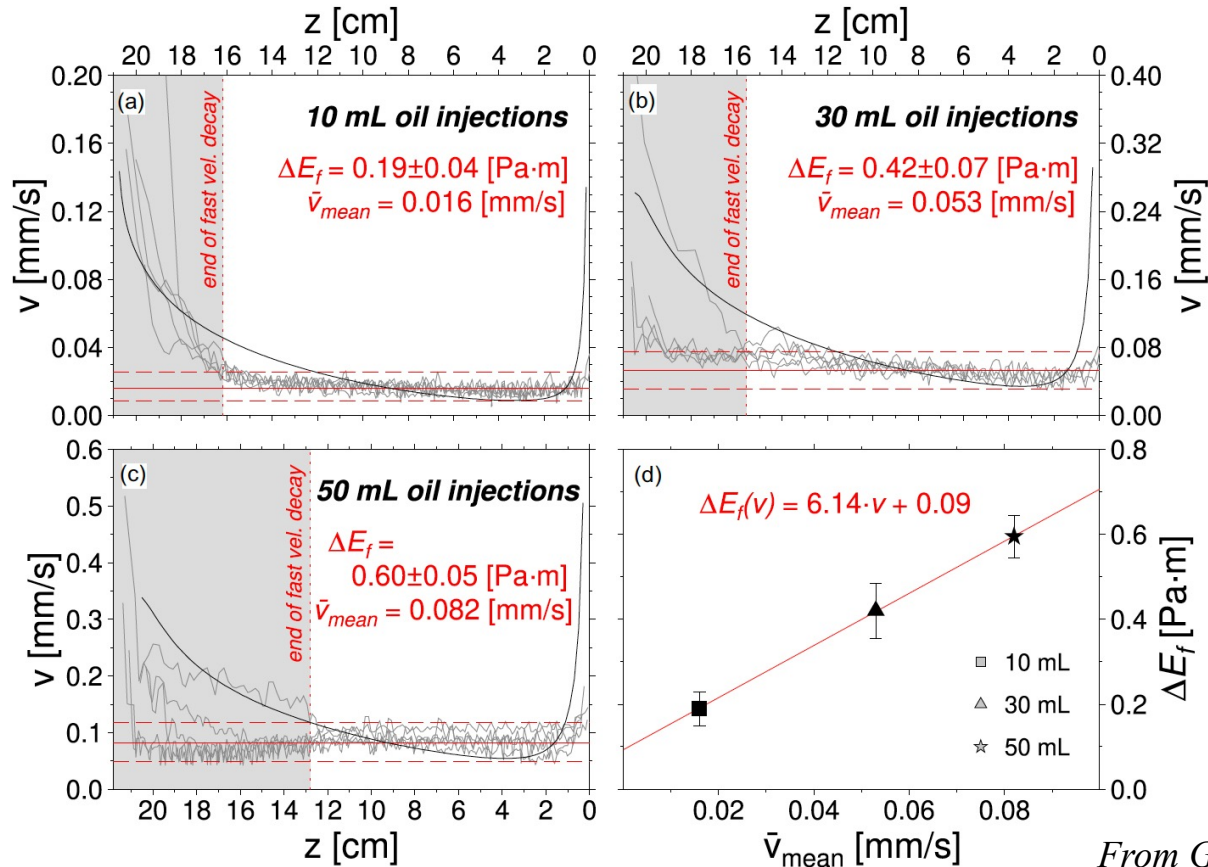


Oil viscosity 0.97 Pas



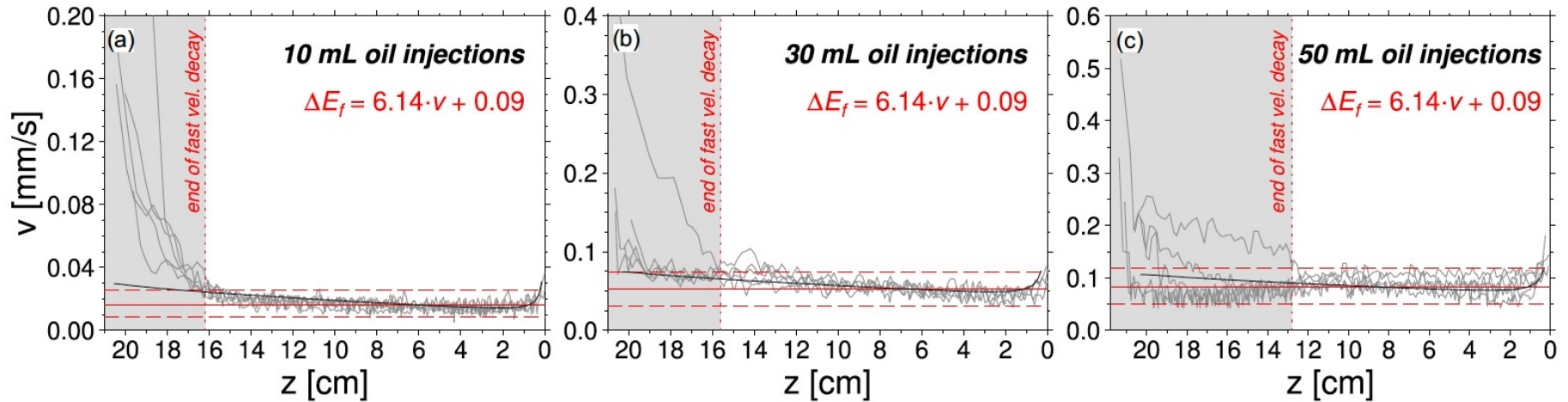
From Furst et al., Volcanica, 2024

Fracture toughness depends on velocity

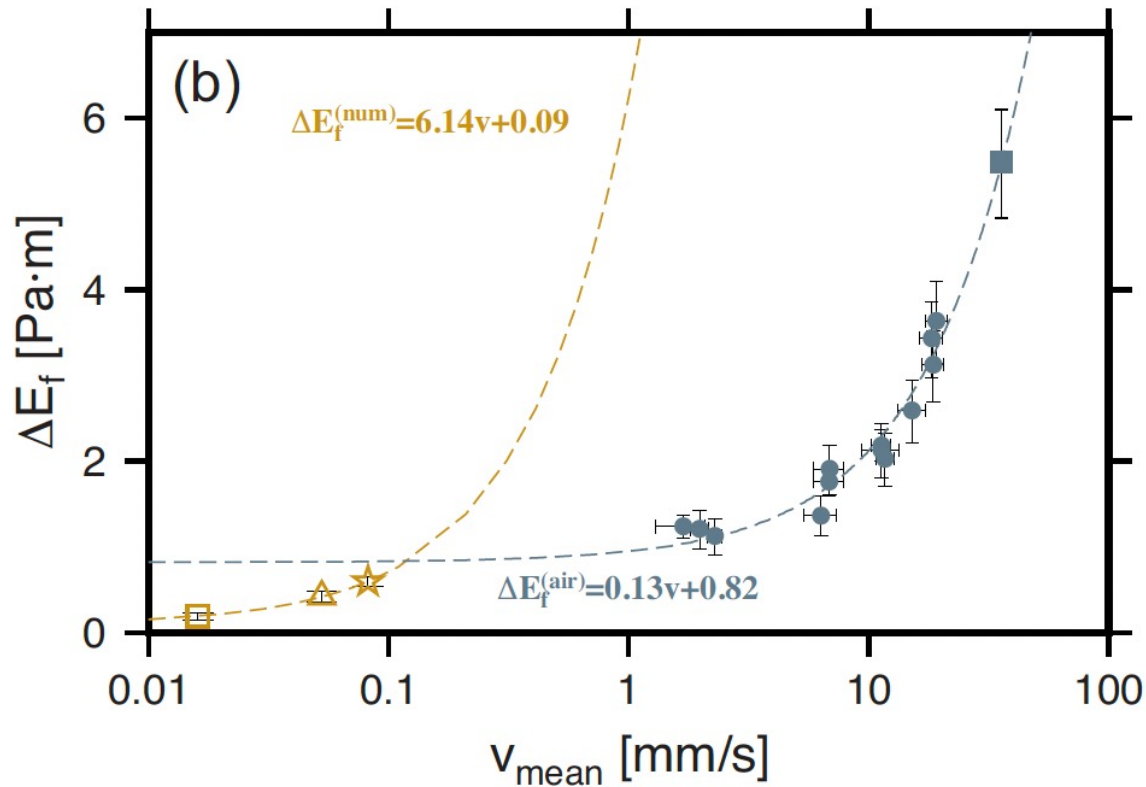


From Gaete et al., GJint, 2025

$Kc(v)$ need to be an input of the numerical simulation



We need to constrain $Kc=f(v)$ for the magma velocities



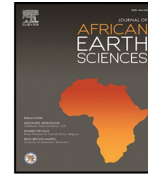
Conclusions on orientation of propagation

**First order of magma path given by the local stress field
at Nyiragongo volcano**

$r < 5\text{km}$ edifice load

$r > 5\text{ km}$ rift extension

Pinel & Mériaux, 2025



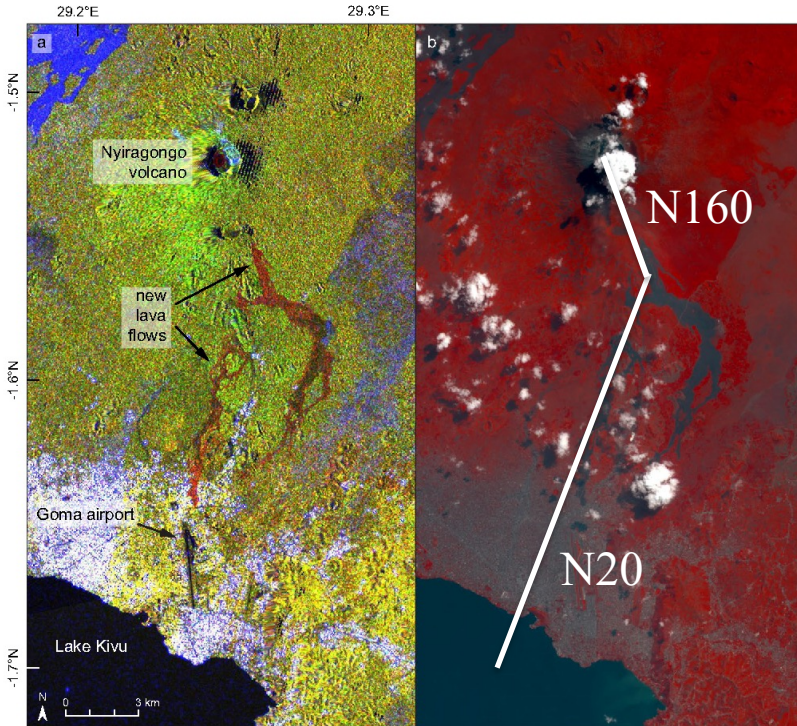
Moving forward:

-We need a better stress field characterization (faults)

-2D models of magma propagation can improve the path prediction and provide information on the velocity but a better characterization of the crustal fracture toughness is needed

Nyiragongo volcano, DRC: 20 km of lateral propagation

1977, Jan. 2002, May 2021



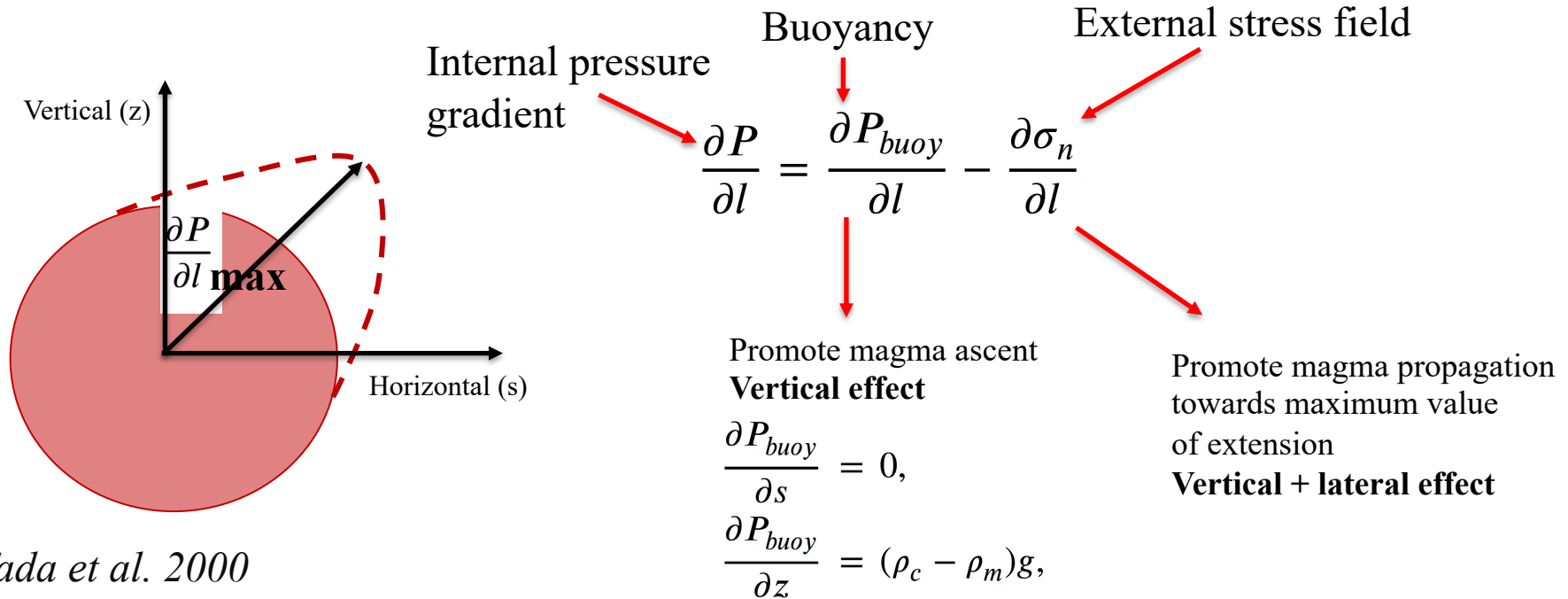
Orientation → magma path
 $r < 5\text{km}$ edifice load
 $r > 5\text{ km}$ rift extension

Direction of propagation
→ lateral versus vertical

From Poland, 2022

Theoretical principle

A penny shaped crack will propagate in a direction corresponding to the **maximum pressure gradient** (quasi static case) → maximum stress intensity factor K_I



Tada et al. 2000

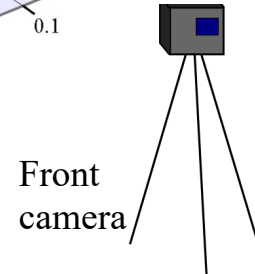
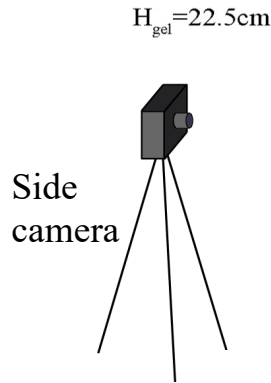
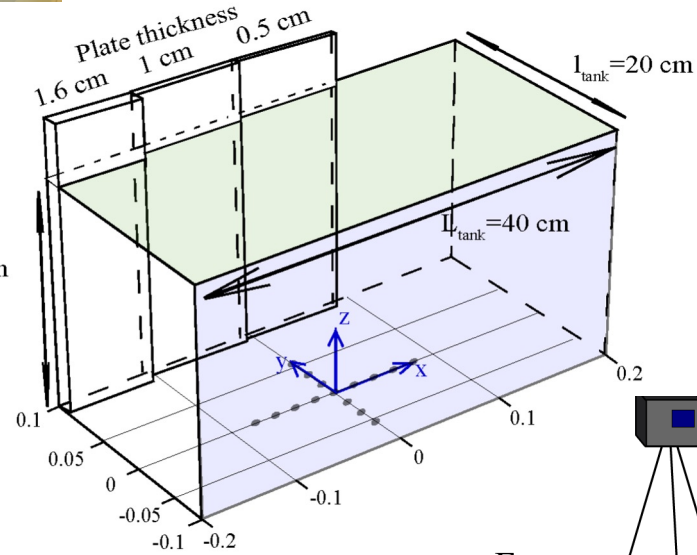
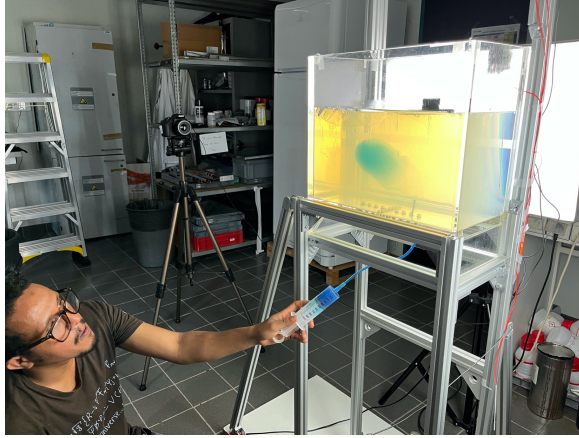
Davis et al., 2021

This approach neglects the potential effects of:

- the dyke non-isotropic shape
- magma flow

→ Need for validation

Experimental validation: Oil injection inside gelatin



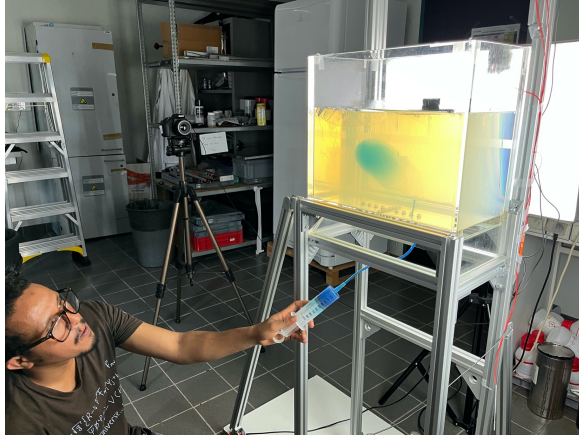
Measured gelatin properties:

- density ($1000 < \rho_{\text{gel}} < 1104 \text{ kg/m}^3$)
- Young's modulus ($1250 < E_{\text{gel}} < 2193 \text{ Pa}$)
- (*Smittarello et al. 2021*)
- ν close to 0.5

Known oil properties:

- density ($\rho_{\text{oil}} = 871 \text{ kg/m}^3$)
- viscosity ($\mu_{\text{oil}} = 1.74 \cdot 10^{-3} \text{ Pas}$)

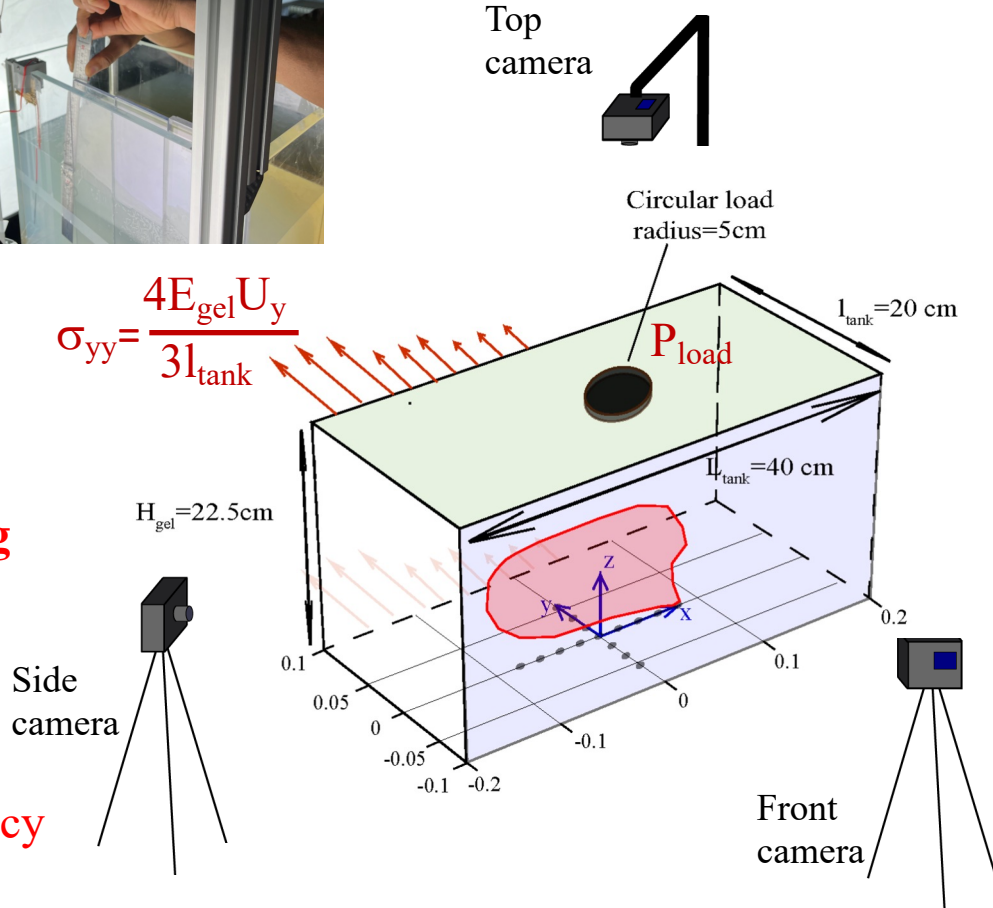
Experimental validation: Oil injection inside gelatin



Extension imposed by removing lateral plates

Varying:

- surface load (P_{load})
- extension amount (σ_{yy})
- density contrast ($\Delta\rho$) $\rightarrow$ buoyancy



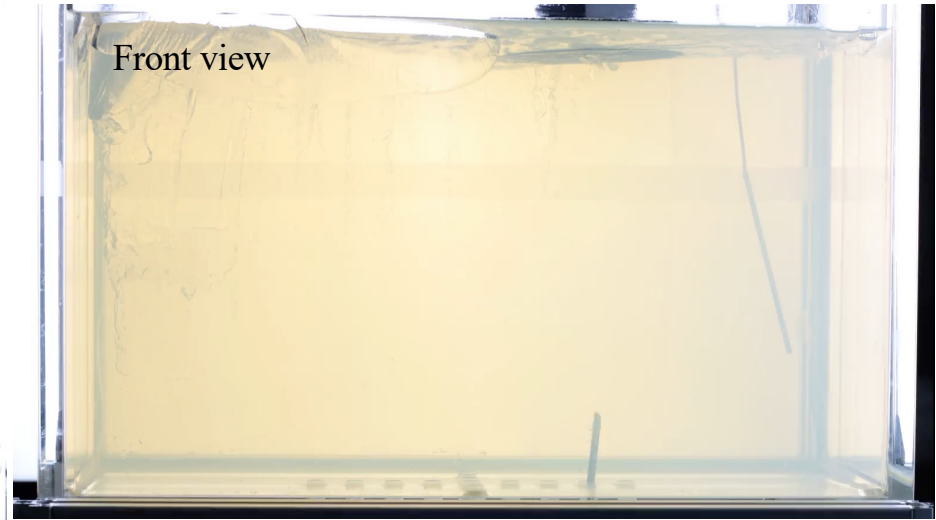
Extension $\sigma_{yy}=234\text{Pa}$, $d\sigma_{yy}/dx=975\text{Pa/m}$

$P_{\text{load}} = 138 \text{ Pa}$

$\Delta\rho=129$, $E= 2193 \text{ Pa}$

Film accelerated *1200

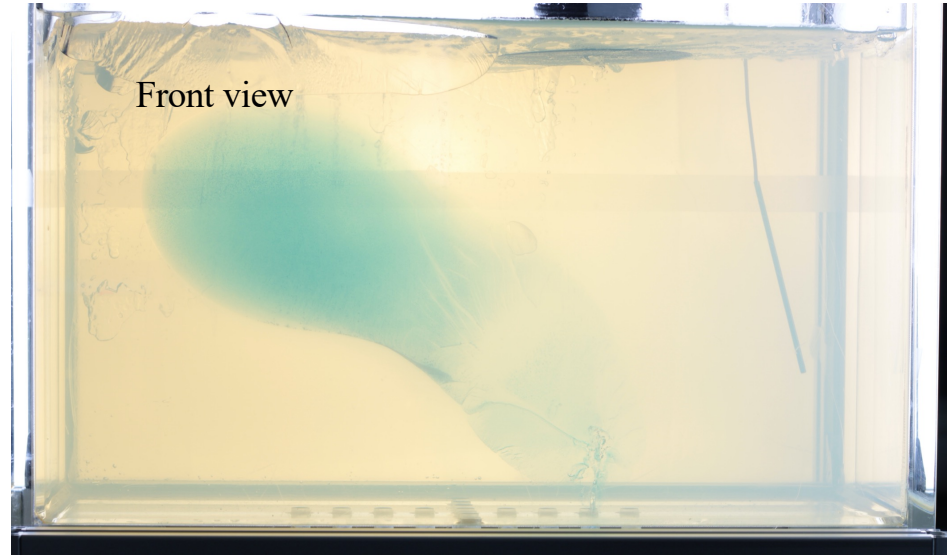
True duration: 36 min



Extension $\sigma_{yy}=234\text{Pa}$, $d\sigma_{yy}/dx=975\text{Pa/m}$

$P_{\text{load}} = 138 \text{ Pa}$

$\Delta\rho=129$, $E= 2193 \text{ Pa}$



Extension $\sigma_{yy}=234\text{Pa}$, $d\sigma_{yy}/dx=975\text{Pa/m}$

$P_{\text{load}} = 138\text{ Pa}$

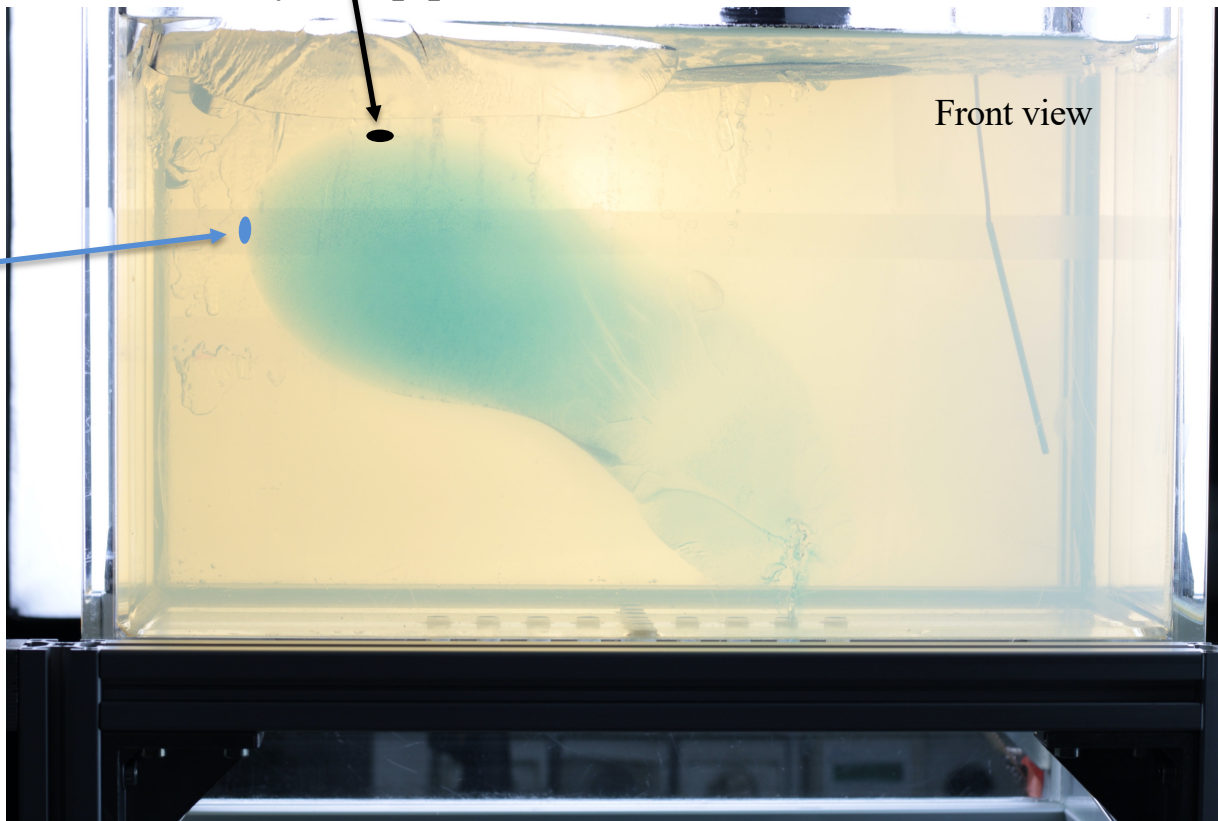
$\Delta\rho=129$, $E= 2193\text{ Pa}$

Surface not reached

Dyke top point

Front view

Dyke lateral
point

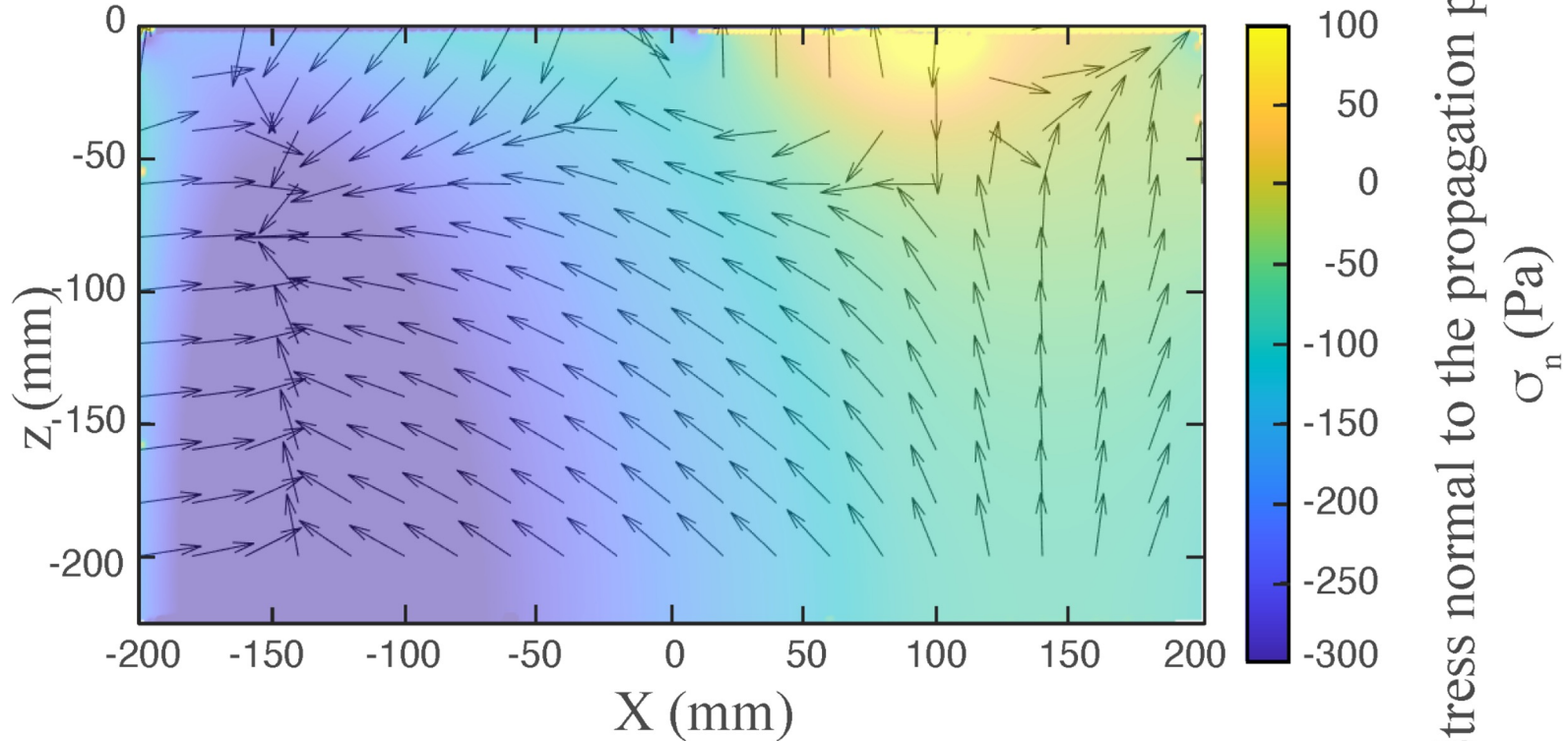


Extension $\sigma_{yy}=234\text{Pa}$, $d\sigma_{yy}/dx=975\text{Pa/m}$

$P_{\text{load}} = 138\text{ Pa}$

$\Delta\rho=129$, $E= 2193\text{ Pa}$

Tank stress field modelled with COMSOL
Mid-plane $Y=0$



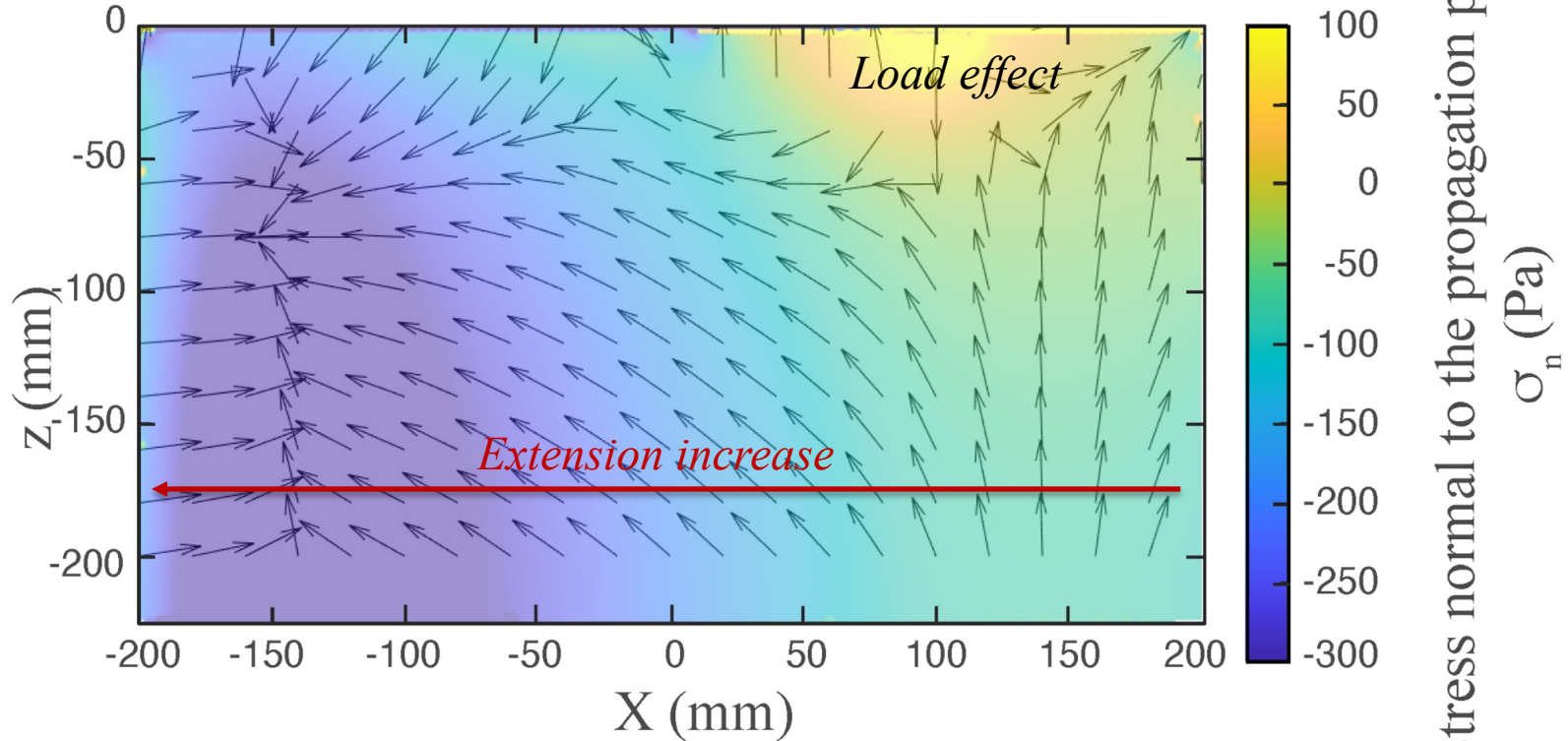
Extension $\sigma_{yy}=234\text{Pa}$, $d\sigma_{yy}/dx=975\text{Pa/m}$

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Tank stress field modelled with COMSOL

Mid-plane $Y=0$

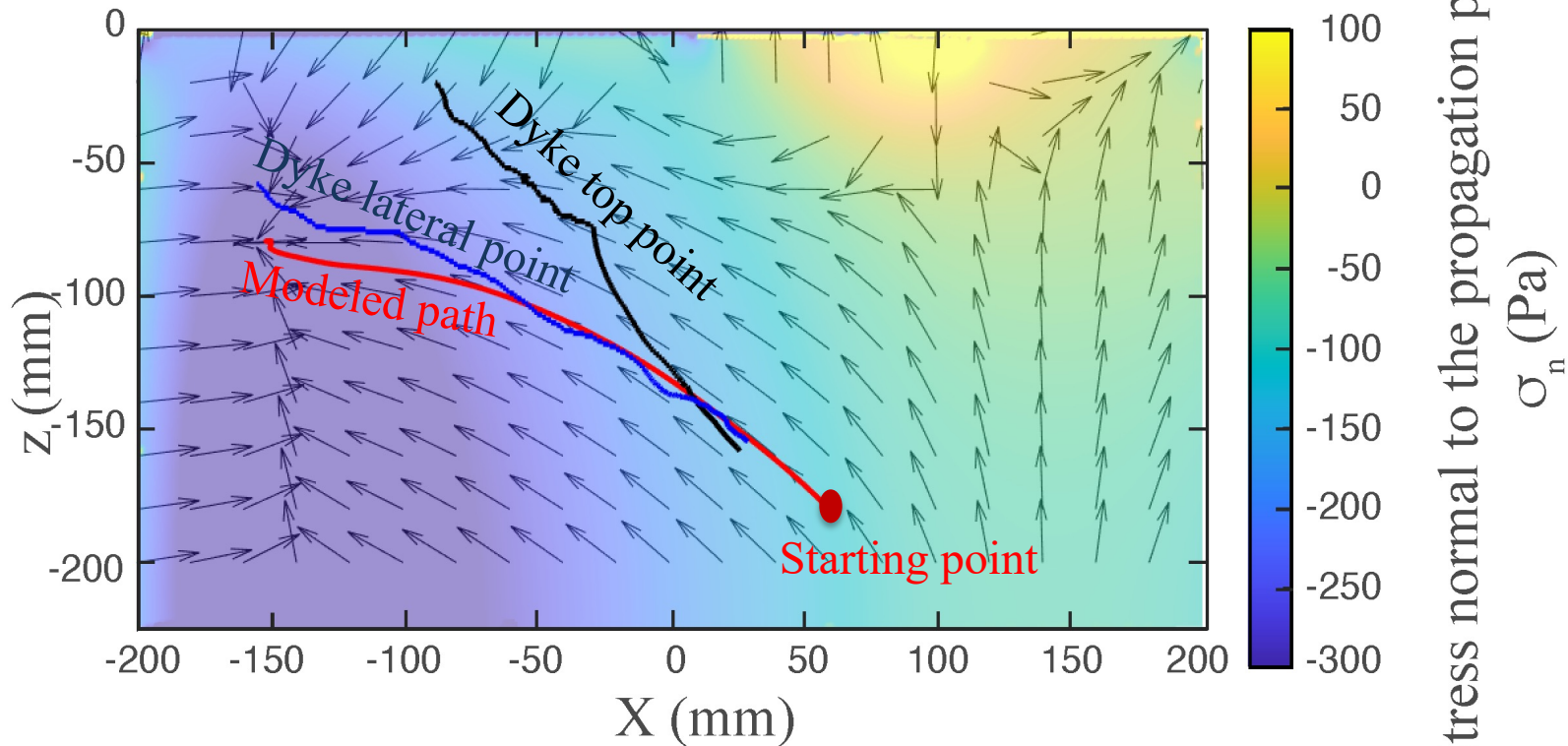


Extension $\sigma_{yy}=234\text{Pa}$, $d\sigma_{yy}/dx=975\text{Pa/m}$

$P_{\text{load}} = 138\text{ Pa}$

$\Delta\rho=129$, $E= 2193\text{ Pa}$

Tank stress field modelled with COMSOL
Mid-plane $Y=0$



Experimental results: 3 other cases

Extension: $d\sigma_{yy}/dx=826\text{Pa/m}$

No load

$\Delta\rho=129\text{ kgm}^{-3}$

Surface not reached

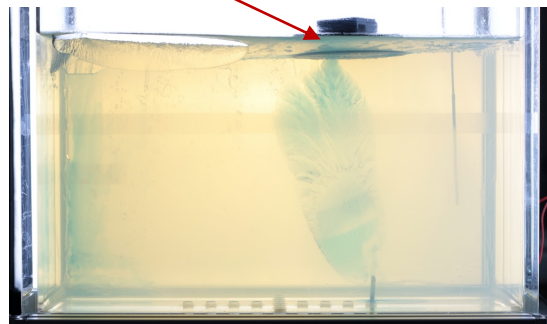


Extension: $d\sigma_{yy}/dx=600\text{Pa/m}$

$P_{\text{load}}=138\text{ Pa}$

$\Delta\rho=233\text{ kgm}^{-3}$

Surface reached

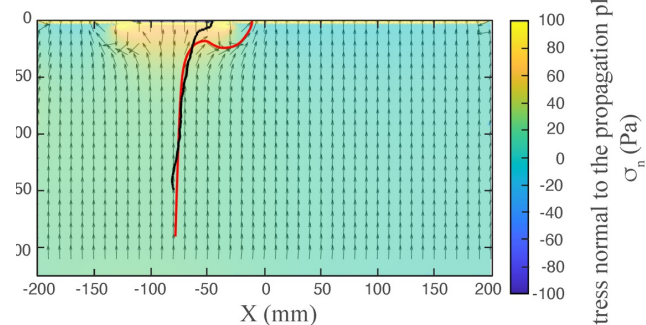
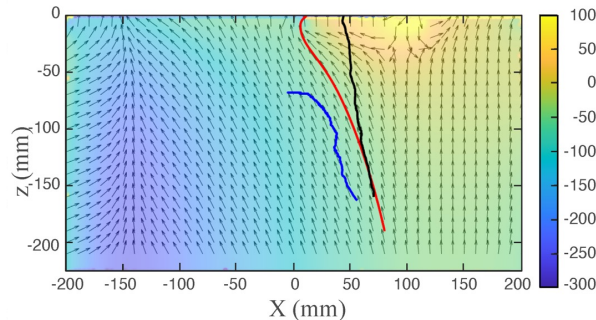
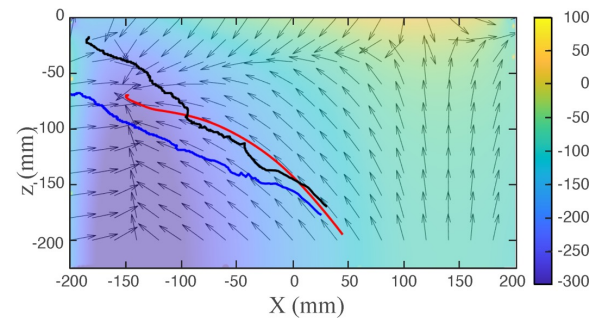
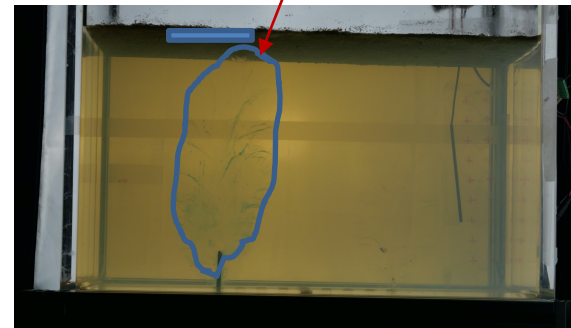


No extension

$P_{\text{load}}=133\text{ Pa}$

$\Delta\rho=131\text{ kgm}^{-3}$

Surface reached

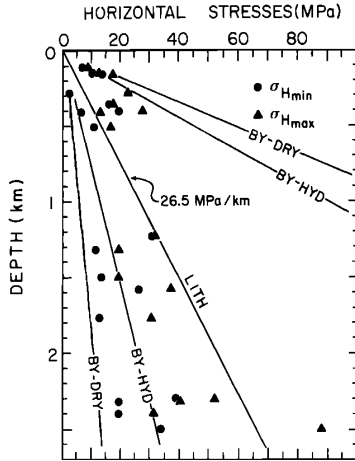


Stress normal to the propagation plane
 σ_n (Pa)

External stress field gradient calculation for Nyiragongo case study

- extension N110°E

A Vertical gradient



(Brace & Kohlstedt., 1980)

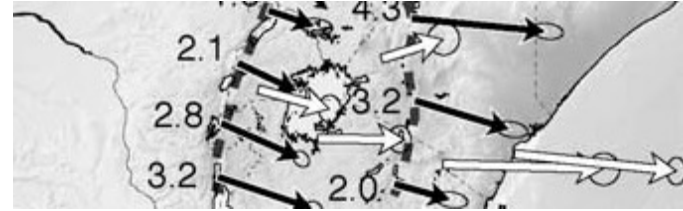
$$\frac{\partial \sigma_n}{\partial z}$$

between 1400Pa/m
and 2500 Pa/m

→ may inhibit vertical ascent

A lateral gradient

estimated from displacement gradient



(Stamps et al., 2008)

Spreading rate locally increasing towards the South
(lateral gradient around 2.2 μmyr⁻¹km⁻¹)

several orders of magnitude smaller

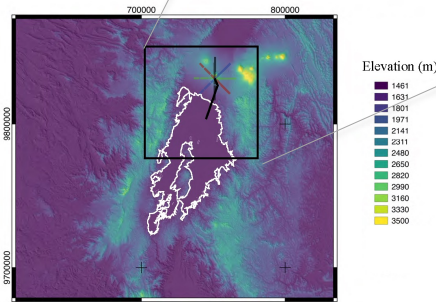
-topography + lake loading effects estimated using a numerical model

Modeling strategy

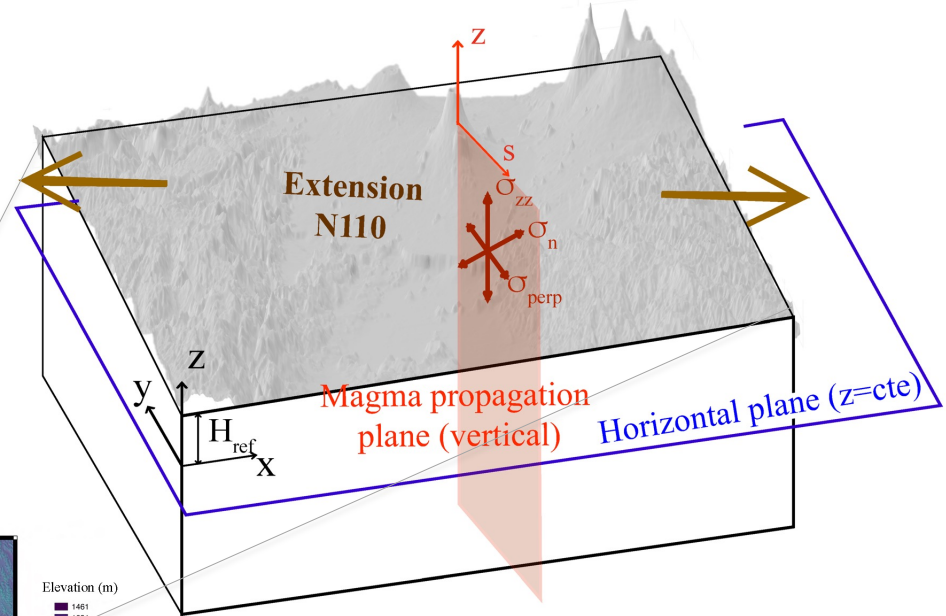
3D stress field calculation:

- Lithostatic reference state
- Extension (N110)
- Topographic load + water load**
applied at $H_{\text{ref}}=1.461$ km (lake level)

σ_n : Stress normal to the vertical
propagation plane



SRTM DEM
+ bathymetry of Lake Kivu
(Lahmeyer, 1998; Ross et al., 2014)



Numerical solution

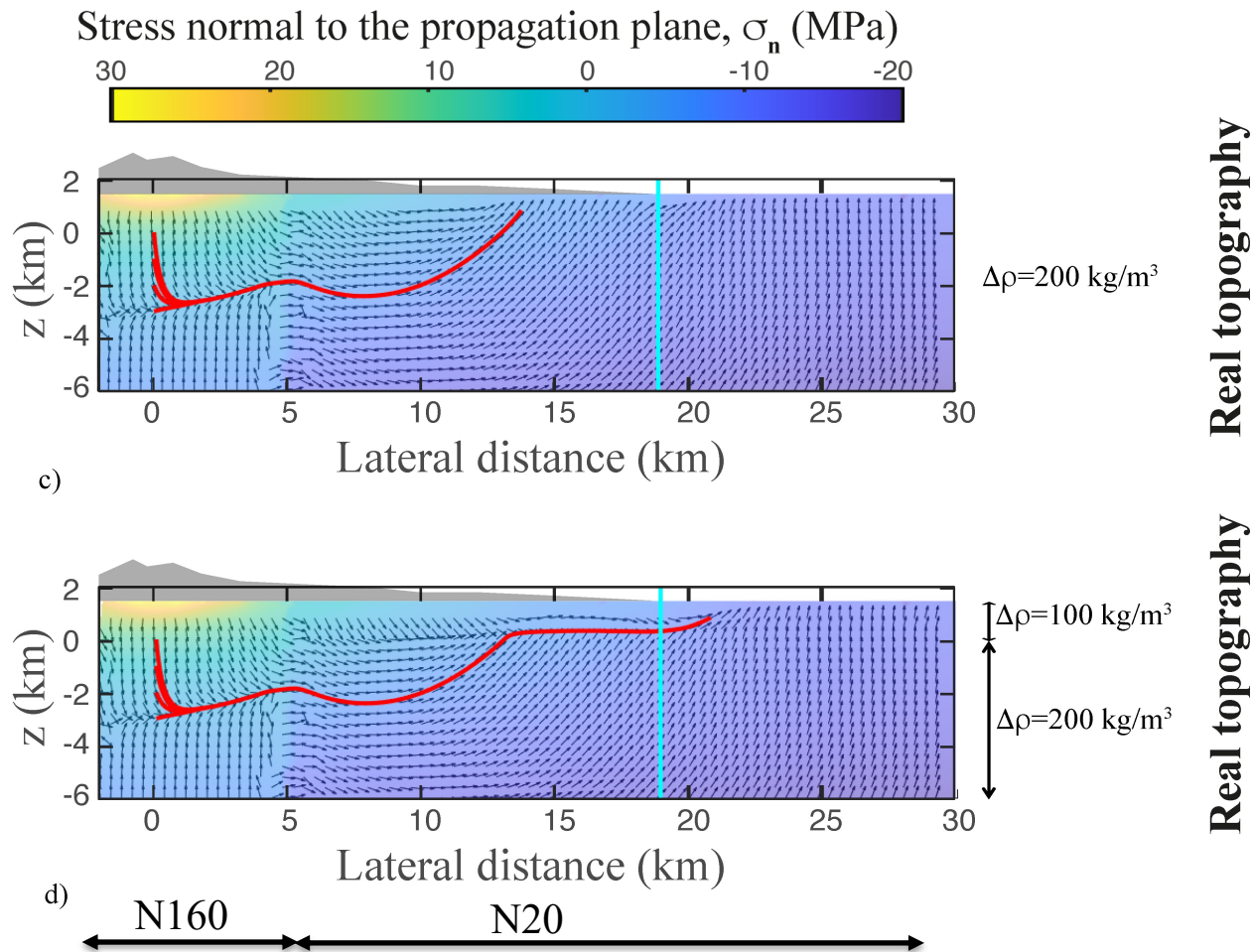
(COMSOL, $80 \times 80 \times 80$ km³ domain)

Zero displacement condition bottom + lateral boundaries

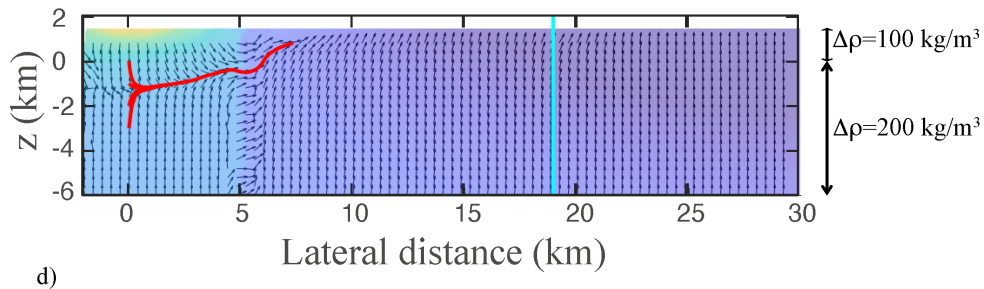
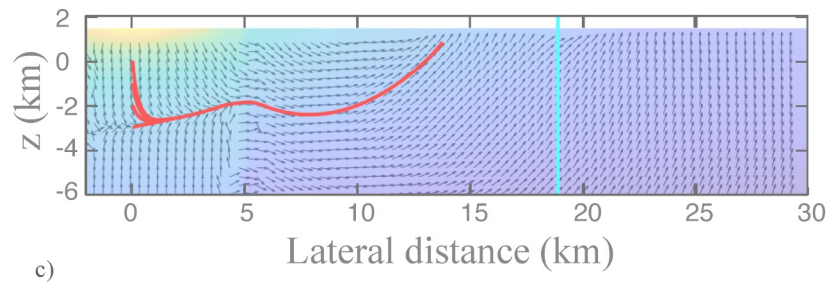
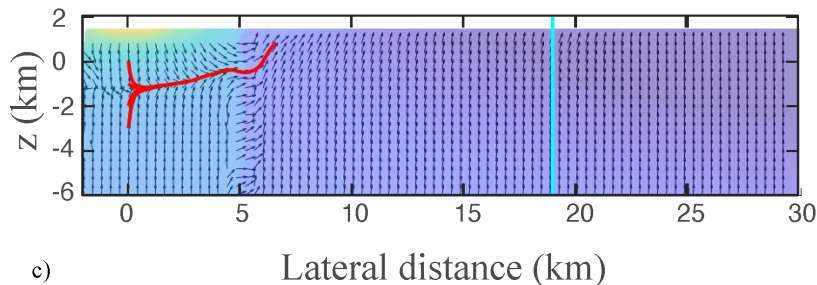
Lateral propagation also requires reduced buoyancy at shallow depth

Extension gradient:

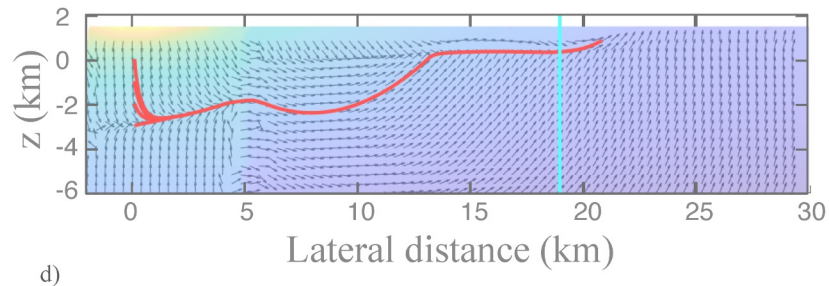
- Vertical: -2500Pa/m
- 15MPa at 6 km depth
- Horizontal: +0.002 Pa/m



The vertical gradient of the extension is really needed



Without vertical gradient



With vertical gradient

Lateral propagation is favored by:

- **depth-dependent extension**
- **reduced magma buoyancy at shallow depth**
- downslope towards Lake Kivu
- slight increase in rift extension towards the south

Conclusions on direction of propagation

Rift zone are prone to magma lateral propagation

- Magma may or may not reach the surface
- If magma does reach the surface, it may be very far away (several tens of km) from its source

This must be taken into account when assessing volcanic hazards and associated risks.

Modeling the gradient of magma overpressure → direction of magma propagation

Validated by analogue modeling

Applied to Nyiragongo volcano

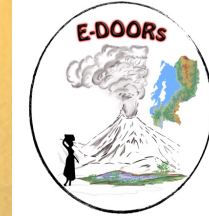
- **deviatoric stress in extension increases with depth in rift zones**
- **magma buoyancy is reduced** at shallow depth around Lake Kivu

General conclusion

Combining analogue experiments and numerical models is a good way to better understand magma propagation

We need :

- **To better characterize the local stress field**
- **To better characterize crustal properties (fracture toughness)**



Thank you for your attention...