Advanced School on Earth System Modelling & Workshop on Climate Change and Regional Impacts over South Asia

- **1. Modeling Ocean Biogeochemistry**
- 2. IITM-ESM and biophysical feedbacks in the Indian Ocean



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What is ocean biogeochemistry?

Biology – micro-scale
Chemistry – organic and inorganic
Geology – interactions with solid Earth
Physical interactions
Air-sea exchange; Particle settling rates;
Advection, diffusion, mixing

Why include biogeochemistry in ocean models?

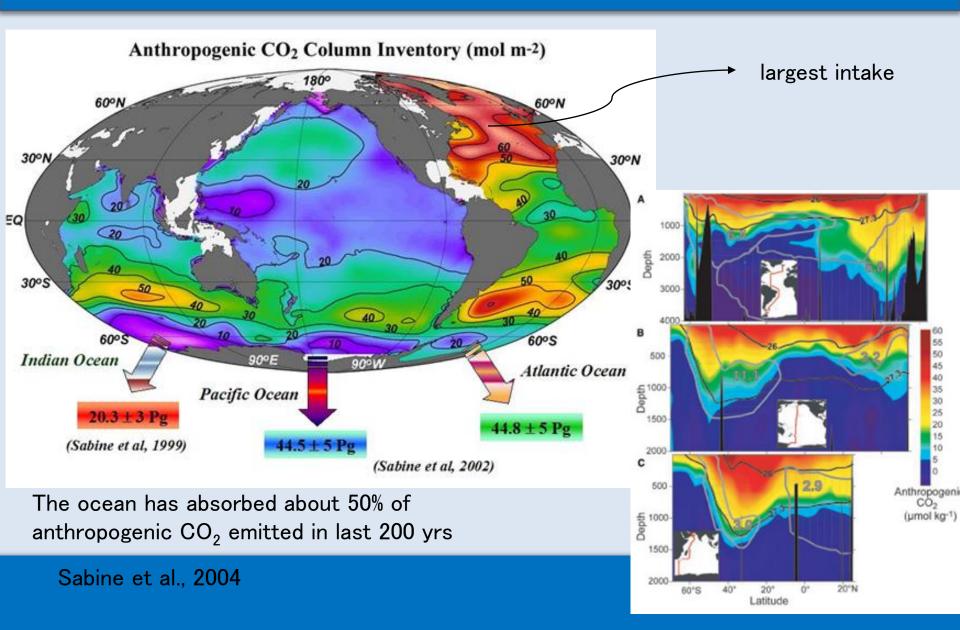
Carbon Cycle

Ocean carbon sink – past, present, future Glacial / interglacial change

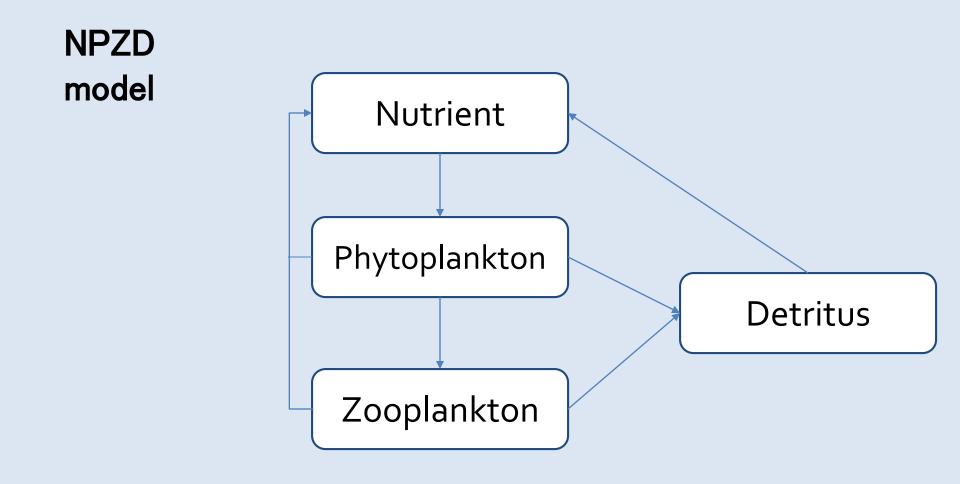
Biophysical feedbacks

Trace gas emissions – Atmospheric chemistry e.g. Dimethyl Sulfide (DMS): CCN, emitted by phytoplankton, theorized climate feedbacks

Why include biogeochemistry in ocean models?



Simple ecosystem model for the lower trophic levels



There are \sim 20,000 of identified species of phytoplankton in 4 major groups

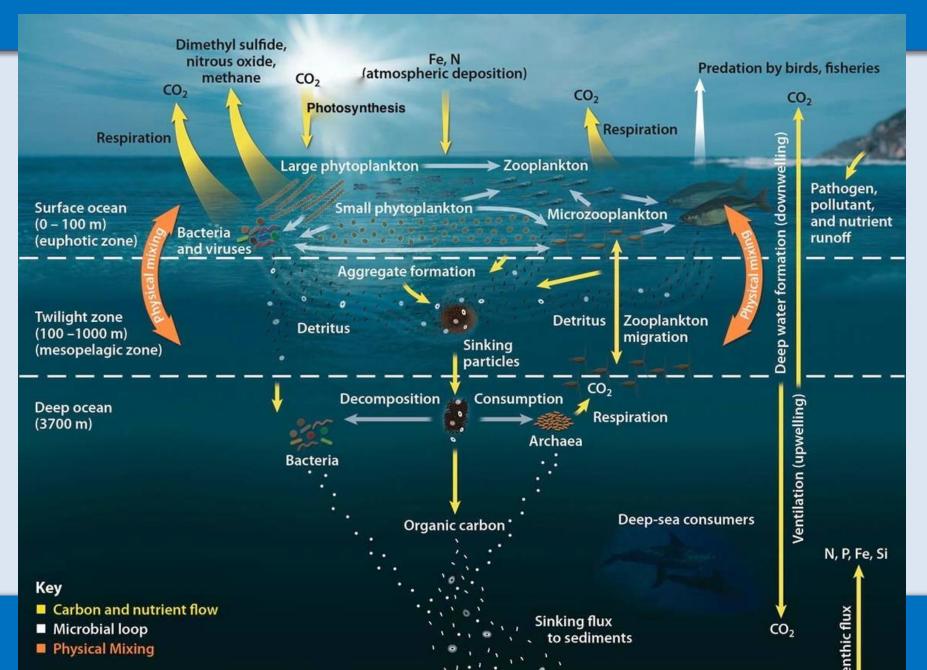
- Picoplankton
- Diatoms (silicate shells)
- Coccolithophorids (carbonate shells)
- Dinoflagellites

Zooplankton – also great variety

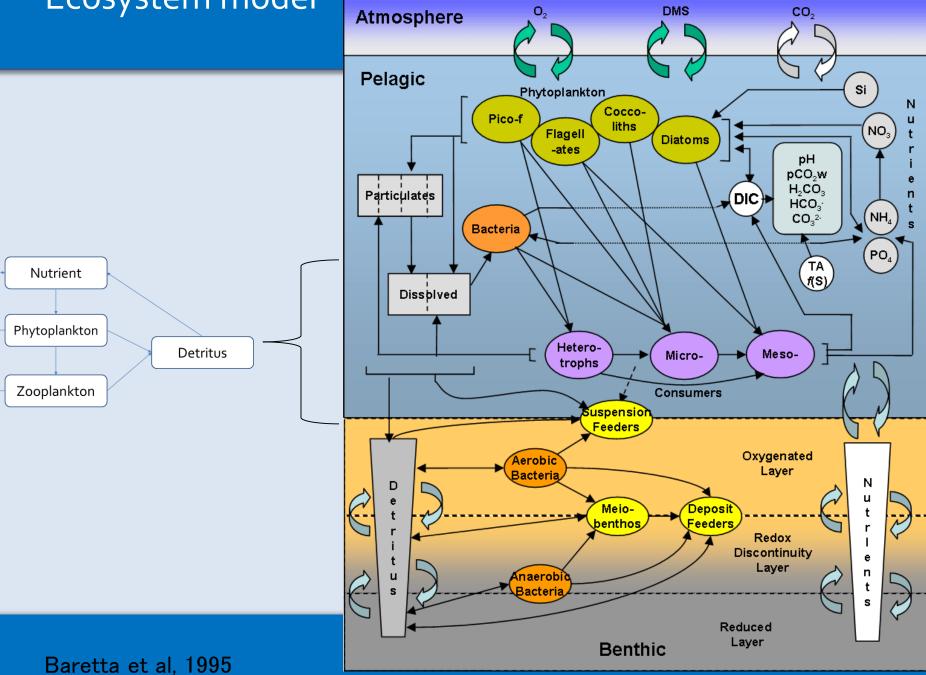
Much variability in key aspects

- Carbon to Nutrient, Carbon to Chlorophyll ratios
- Sinking velocities
- Growth rates, mortality rates, etc.

Biological Pump



Ecosystem model



Ecosystem model

$$\frac{\partial Chl}{\partial t} = (\mu - L) * Chl - (s - w) \frac{\partial Chl}{\partial z} + \frac{\partial \left[K(\partial Chl/\partial z)\right]}{\partial z}$$
(1)

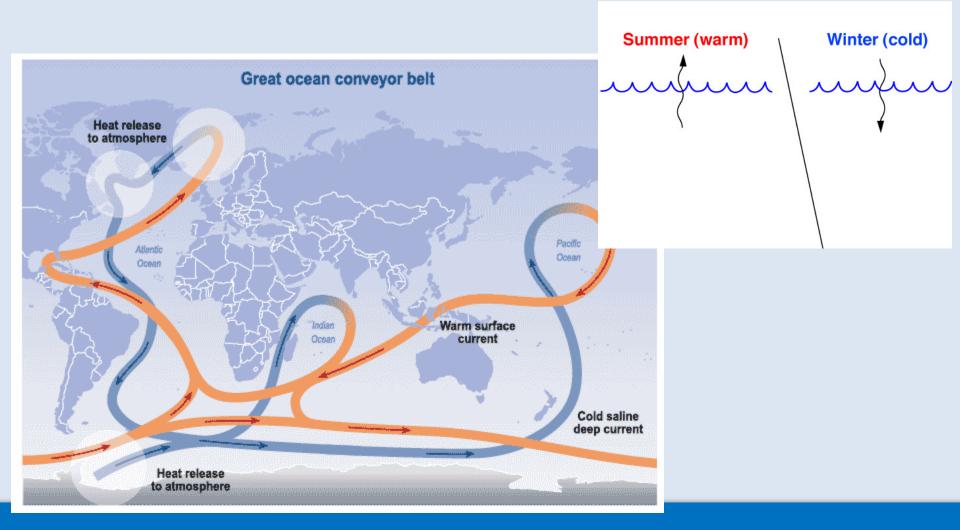
$$\frac{\partial PON}{\partial t} = v * Chl - (s - w) \frac{\partial PON}{\partial z} + \frac{\partial \left[K(\partial PON / \partial z)\right]}{\partial z} - L * PON$$
(2)

$$\frac{\partial N}{\partial t} = -v * Chl + w \frac{\partial N}{\partial z} + \frac{\partial \left[K(\partial N / \partial z)\right]}{\partial z} + r * L * PON$$
(3)

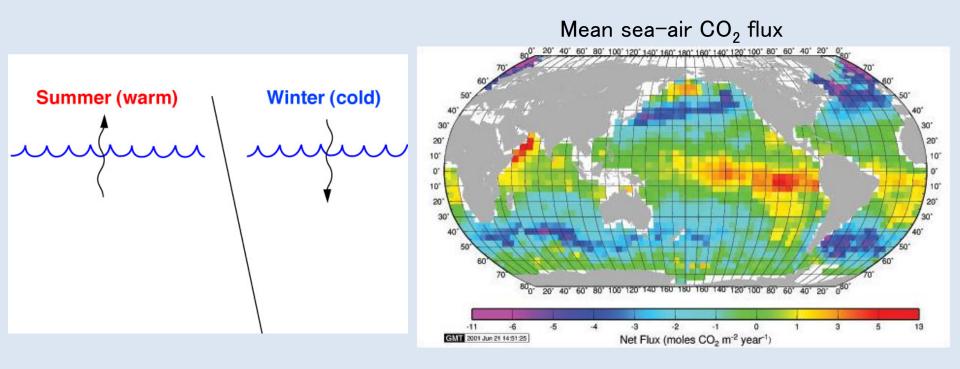
where

- μ : specific growth rate, a function of the internal nutrient concentration and PAR
- L : specific loss rate
- s: phytoplankton sinking rate
- w: upwelling velocity
- K: turbulent vertical eddy-diffusion coefficient
- v: specific uptake-rate
- r: recycling efficiency

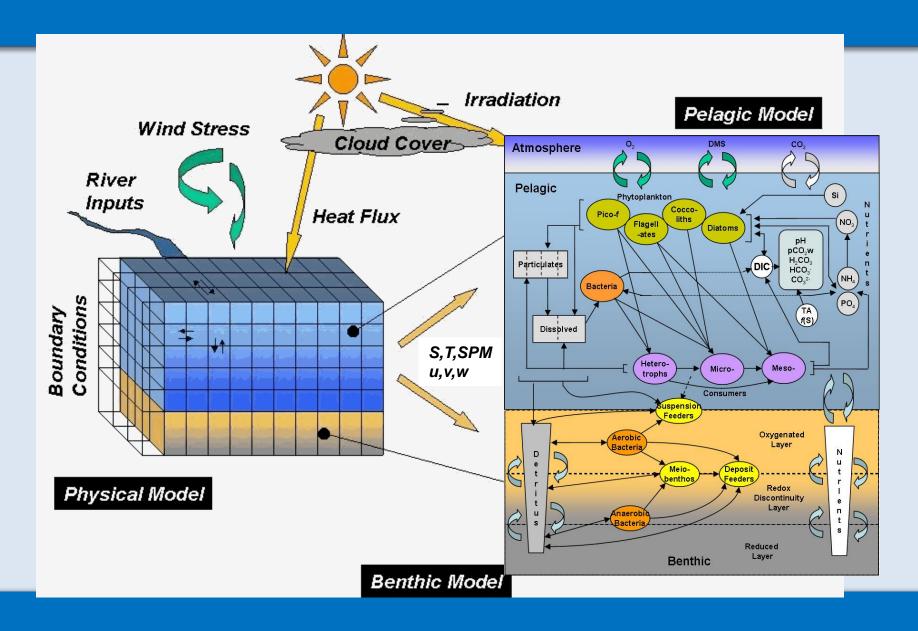
Solubility pump: dissolved, inorganic carbon 1. thermohaline circulation 2. solubility = inverse function of seawater temperature



Temperature influence on Carbon fluxes



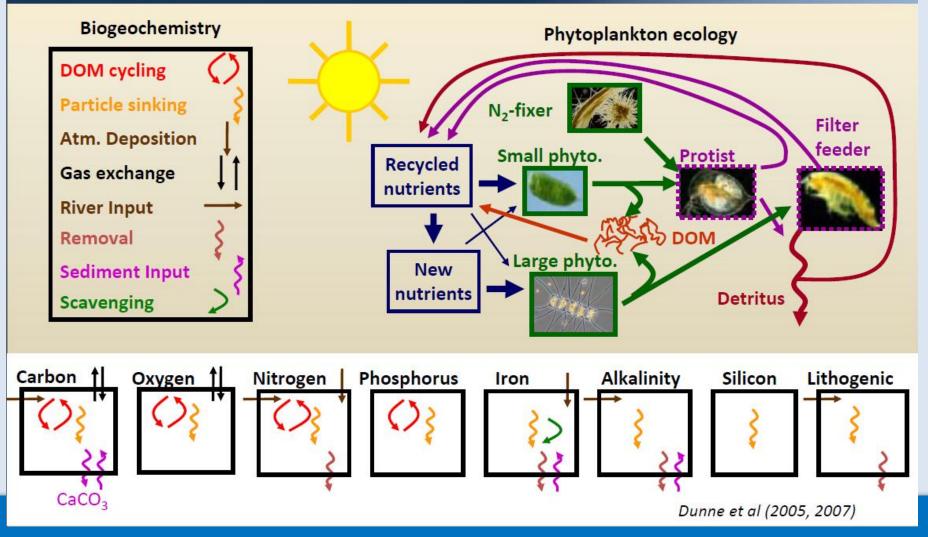
Conceptual coupled Physical-Ecosystem model



Baretta et al, 1995

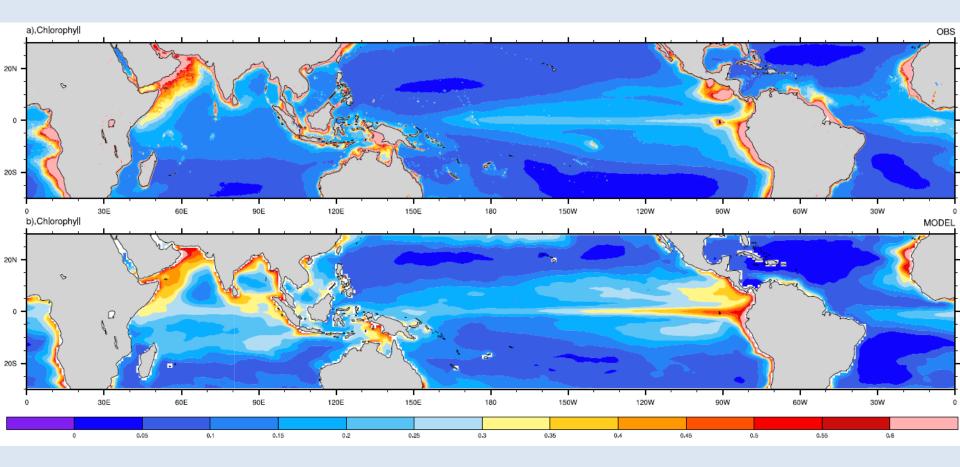
TOPAZ in **IITM-ESM**

Tracers Of Phytoplankton with Allometric Zooplankton (TOPAZ) simulates the mechanisms that control the ocean carbon cycle

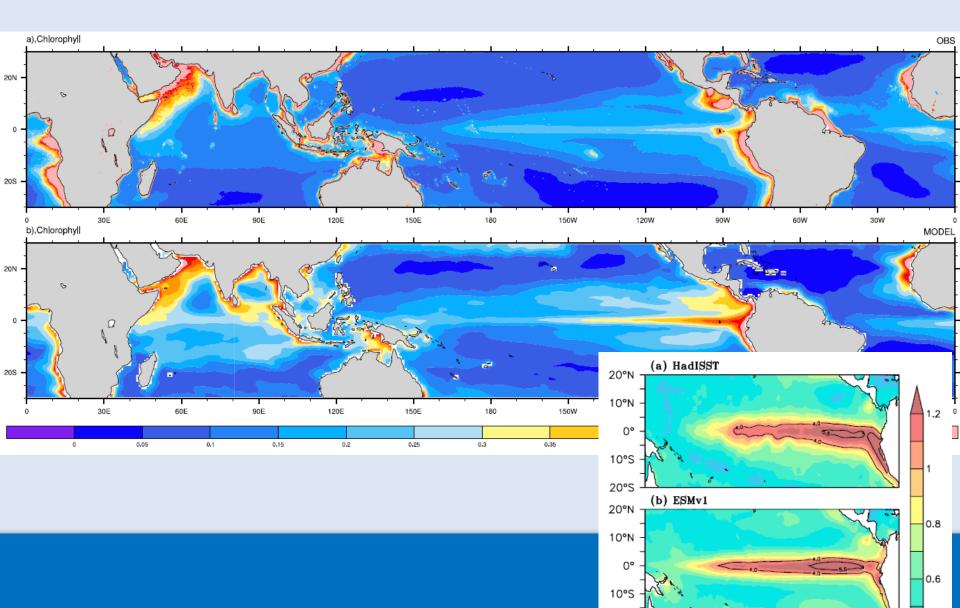


Dunne et al., 2005, 2007

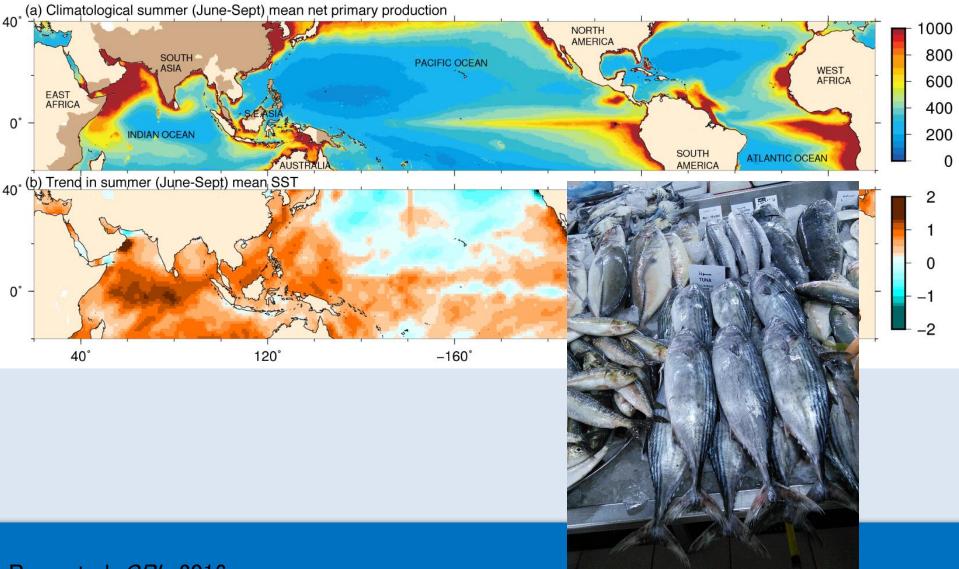
TOPAZ in IITM-ESM Summer Chlorophyll simulations



TOPAZ in IITM-ESM Biophysical feedbacks important



Warming – Marine Primary Production western Indian Ocean is a highly productive region...



Roxy et al. GRL, 2016

Tuna

Warming – Marine Primary Production Earlier studies suggest an increase in WIO NPP

Studies have suggested an decrease in chlorophyll and marine primary production in the tropical oceans, due to rising SSTs.

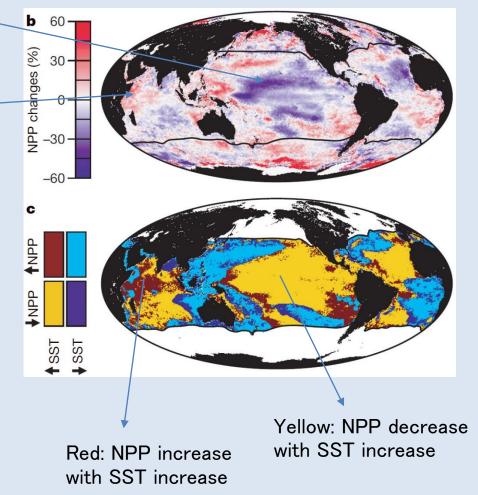
However, the chlorophyll changes in the WIO suggested an increasing trend.

Goes et al. 2005: Increase of more than 350% in summer plankton biomass in the WIO, driven by strengthening monsoon winds (1997–2004).

Gregg et al. 2005: Second largest increase (37%) in Chl among the open oceans, in WIO (1998-2003).

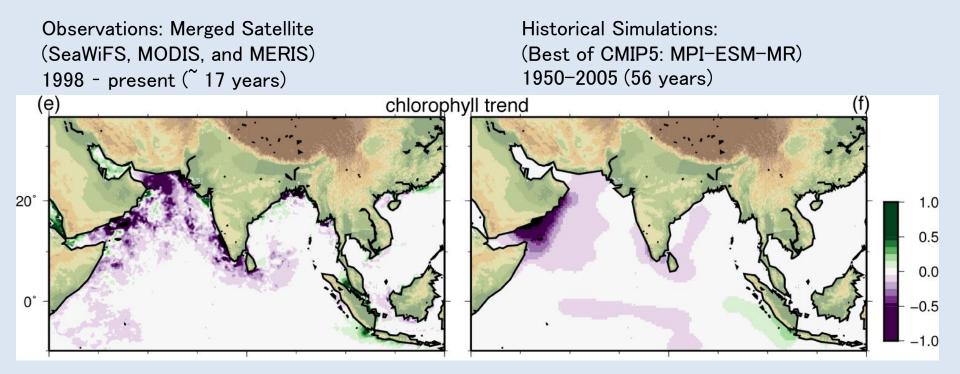
50-60 years of data is required to detect a trend above the natural variability. For tropical regions (WIO) it can be shorter (20-30 years) (Henson et al. 2010, Beaulieu et al. 2013)

Behrenfeld et al. Nature, 2006



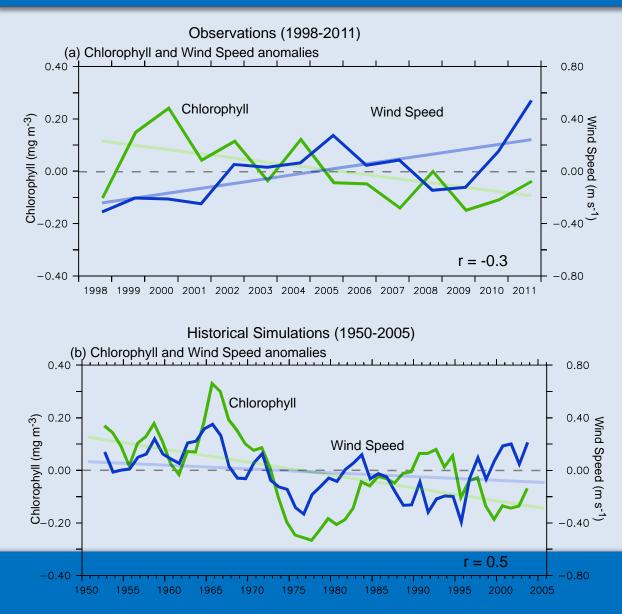
Goes et al. *Science*, 2005; Gregg and Rousseaux, *JGR*, 2005 Behrenfeld et al. *Nature*, 2006

Reduction in Marine Primary Production Chlorophyll trends in observations and simulations



20-30% reduction in marine primary productivity over the western Indian Ocean

Long-term changes in winds over WIO are minor - Minimal role on the chlorophyll trends

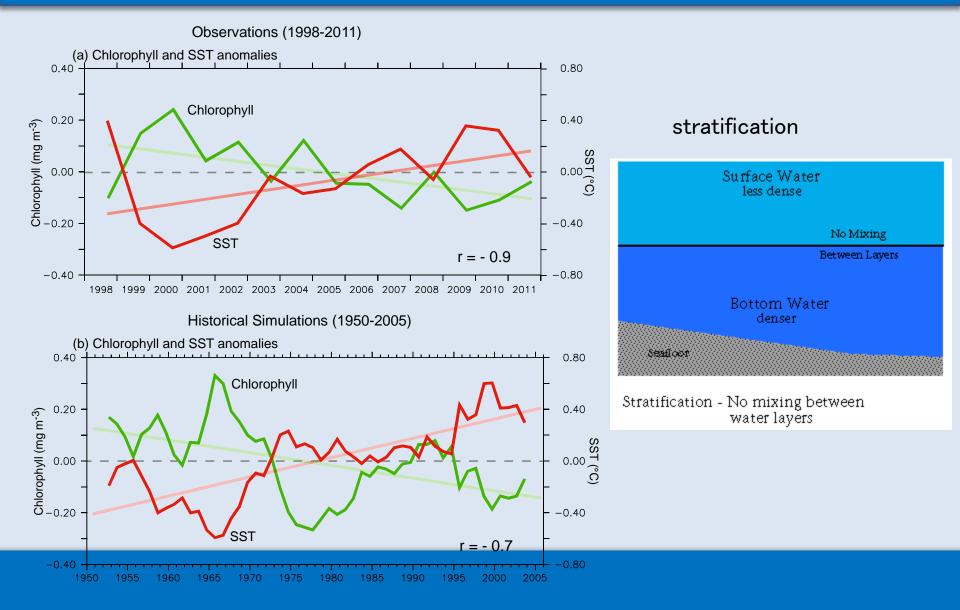


Wind speed anomalies over the western Indian Ocean indicate an increase in wind speed in the recent two decades

However, the long-term changes over the same region is only about 0.2 m s⁻¹ - changes which are minor compared to an SST trend of 0.6C during the same period.

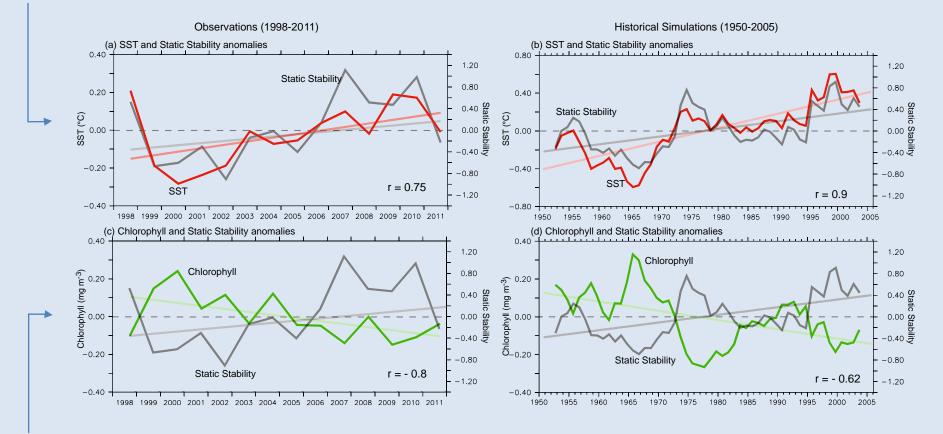
Correlation indicates a a minimal role of the changing winds

Warming stratifies the ocean - and suppresses the mixing of nutrients from the subsurface, reducing chlorophyll



Warming stratifies the ocean - and suppresses the mixing of nutrients from the subsurface, reducing chlorophyll

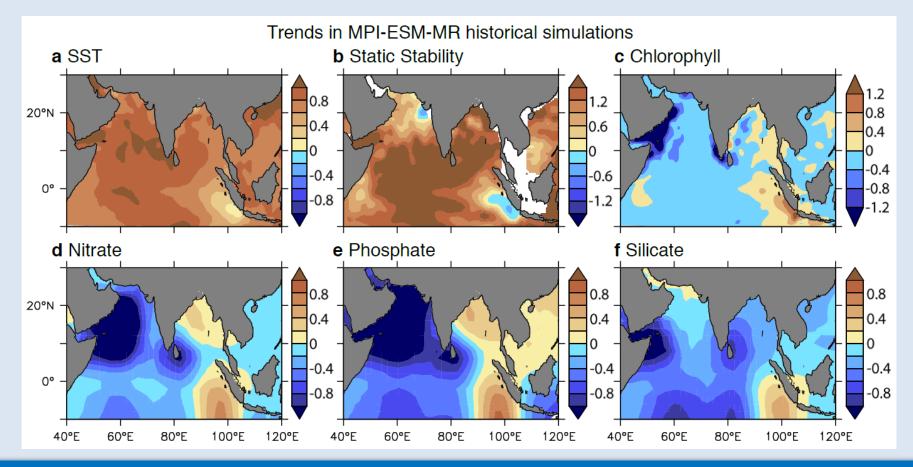




Stratification highly correlated to the reduction in Chlorophyll

Warming stratifies the ocean - and suppresses the mixing of nutrients from the subsurface, reducing chlorophyll

Nitrate, Silicate and Phosphate shows significant reduction over the same region where chlorophyll trend is negative



Earth System Model for South Asia for Future Climate Projections

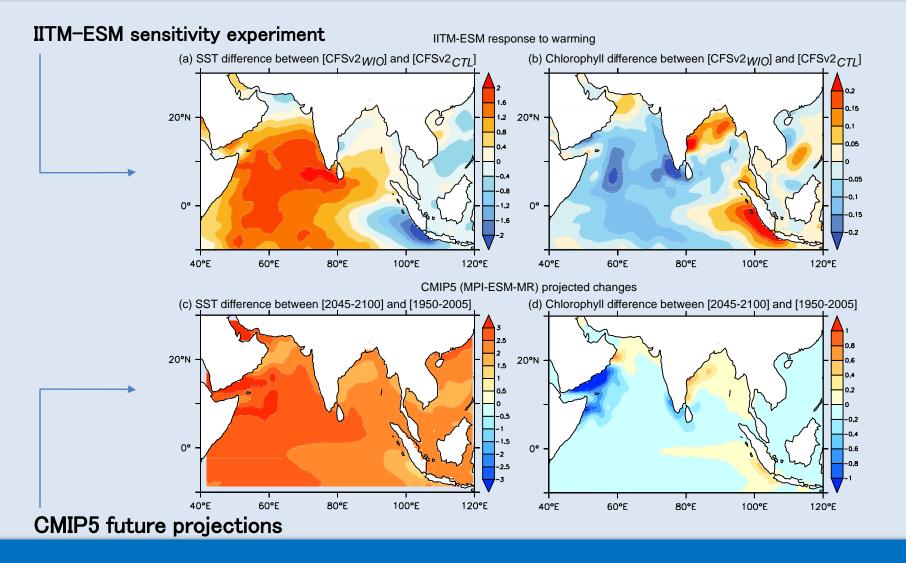
THE IITM EARTH SYSTEM MODEL Transformation of a Seasonal Prediction Model to a Long-Term Climate Model

by P. Swapna, M. K. Roxy, K. Aparna, K. Kulkarni, A. G. Prajeesh, K. Ashok, R. Krishnan, S. Moorthi, A. Kumar, and B. N. Goswami

This work documents the fidelity of the newly developed Indian Institute of Tropical Meteorology climate model simulations and demonstrates its suitability to address the climate variability and change issues relevant to the South Asian monsoon.

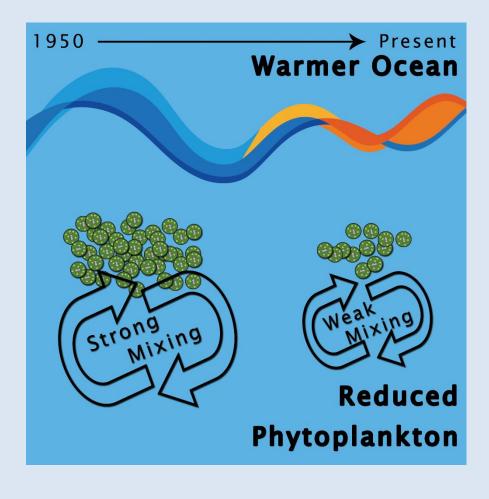
Swapna et al. BAMS, 2015

IITM – Earth System Modelresponse to western Indian Ocean warming



Roxy et al. GRL, 2016

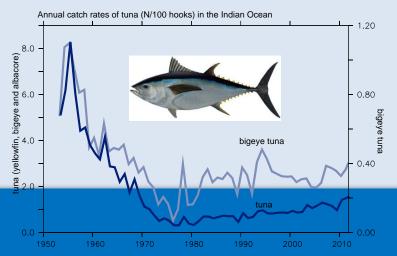
Warming Ocean, Reduced Marine Productivity



Future?

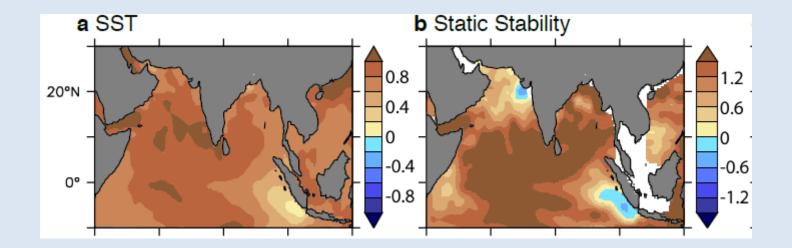
CMIP5 future projections suggest further warming of the Indian Ocean. Will the phytoplankton decrease further?

Is Indian Ocean turning into an ecological desert? Along with the stress from fisheries industries... reduced plankton might increase the fish stress



Roxy et al. GRL, 2016

Missing links - asymmetry in the warming



Identify and separate dynamics/processes leading to surface and subsurface warming

