# **MAGMA RESERVOIRS:**

# EMPLACEMENT AND POST-EMPLACEMENT DYNAMICS



Bulk composition of continental crust dictates that:

- mafic magmas differentiate (to produce evolved compositions),
- mafic cumulates founder or sink (out of the crust).



Magma ascent is buoyancy driven.







#### Rum intrusion, Scotland

| Intrusion    | ho Mafics (kg m <sup>-3</sup> ) | ho Host Rocks (kg m <sup>-3</sup> ) | ho Magma (kg m <sup>-3</sup> ) |
|--------------|---------------------------------|-------------------------------------|--------------------------------|
| Ardnamurchan |                                 |                                     |                                |
| Rum          | 3100-3200                       | 2500-2650                           | 2640                           |
| Sept-Iles    | 3000-3200                       | 2760-2820                           | 2680                           |
| Bushveld     | 3000-3200                       | 2700-2870                           | 2780                           |







### Ardnamurchan, Scotland



Sagging : evidence for dense material in the intrusion (not the initial magma)

Rheological behaviour of crustal rocks: brittle and power-law creep. Stresses in post-crystallization phase: imposed by mafic cumulates.

 $\sigma \approx \Delta \rho g h$ 

 $\Delta \rho$  = density contrast between intrusion and host rocks 300-400 kg m<sup>-3</sup>. h = intrusion thickness 3-8 km. Effective viscosity  $10^{15} - 10^{20}$  Pa s (strain rates >  $10^{-13}$  s<sup>-1</sup>).





| Intrusion    | Volume<br>(km <sup>3</sup> ) | Aspect ratio |
|--------------|------------------------------|--------------|
| Ardnamurchan | 30                           | 1            |
| Rum          | $5.4  10^2$                  | 0.5          |
| Sept-Iles    | $2.8  10^4$                  | 0.2          |
| Bushveld     | $10^{6}$                     | 0.04         |









Driving = buoyancy.

Resisting = viscous shear at top and bottom of intrusion.

Flow of radius  $R^*$ , thickness  $H^*$ .

For flow over radial distance  $R^*$ , kinematic boundary layer extends to distance  $\sim R^*$ . Characteristic velocity (spreading rate)  $U_S$ .

Strain rate  $\sim U_S/R^*$ .



Bulk horizontal momentum balance:

$$(\rho_T g' H^*) H^* R^* \sim \left[ (\mu_- + \mu_+) \frac{U_s}{R^*} \right] R^{*2}$$

Add volume conservation.

$$U_s \sim \frac{\rho_T g' H^{*2}}{\mu_- + \mu_+}$$

$$R(t) \sim \left(\frac{\rho_T g' V^2}{\mu_+ + \mu_-}\right)^{1/5} t^{1/5}$$
$$h(t) \sim \left(\frac{(\mu_- + \mu_+) V^{1/2}}{\rho_T g'}\right)^{2/5} t^{-2/5}$$







End of the spreading phase such that cooling proceeds faster than spreading.

Cooling rate ~ 
$$\kappa/H$$
.  
Spreading rate ~  $\frac{dR}{dt} \sim \left(\frac{\rho_T g' V^2}{\mu_+ + \mu_-}\right)^{1/5} t^{-4/5}$ 

Thus, critical time for buoyancy reversal:

$$t_c \sim \left(\frac{(\mu_- + \mu_+)V^{1/2}}{\rho_T g' \kappa^{5/4}}\right)^{4/9}$$

. . .

At  $t = t_c$ ,  $H = H_c$  and  $R = R_c$ , corresponding to critical aspect ratio  $\alpha_c = H_c/R_C$ :

$$\alpha_c \sim \frac{H_c}{R_c} \sim \left(\frac{(\mu_- + \mu_+)\kappa}{\rho_T g' V}\right)^{1/3} = Ai^{-1/3}$$

Dimensionless number Ai:

$$Ai = \frac{\rho_T g' V}{(\mu_- + \mu_+)\kappa}.$$



 $\alpha_c = (2.8 \pm 0.10) A i^{-1/3}.$ 

 $\alpha_c$  decreases with increasing volume V\*.

Dimensionless number Ai = ratio between two velocity scales. Spreading rate versus and cooling rate (diffusion controlled). Cooling rate.

$$U_d = \frac{\kappa}{H^*}$$

Set  $H^* \sim V^{1/3}$  and  $R^* \sim V^{1/3}$ :

$$\frac{U_s}{U_d} = \frac{\rho_T g' V}{(\mu_- + \mu_+)\kappa} = Ai$$

Large Ai: spreading faster than cooling = intrusion extends to large distances. Small Ai: cooling faster than spreading = intrusion does not spread.







Sagging ? inward dips increase towards axial zone: not consistent with sagging (flexural behaviour) suggests incipient foundering





Sequence of increasing deformation with increasing volume. Sagging – funnel structure – sinking (?)





Sequence of increasing deformation with increasing volume. Complex shapes for small aspect-ratio intrusions.







# Teardrop regime



"Jellyfish" regime

t=710s











## Annular regime











t=870 s









#### Idealized cross-section







## Sinking regime:

very small residual volume at original emplacement depth



### Residual intrusion regime: large volume left at original emplacement

level

BUT

intrusion

strongly deformed























