

# Shallow salinity stratification and air-Sea interaction in the Bay of Bengal

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ASIRI



ICTP June 2016







**Dipanjan Chaudhuri      Sree Lekha**

***“ .... truth flourishes where the student’s light has shone,  
And there alone.”***

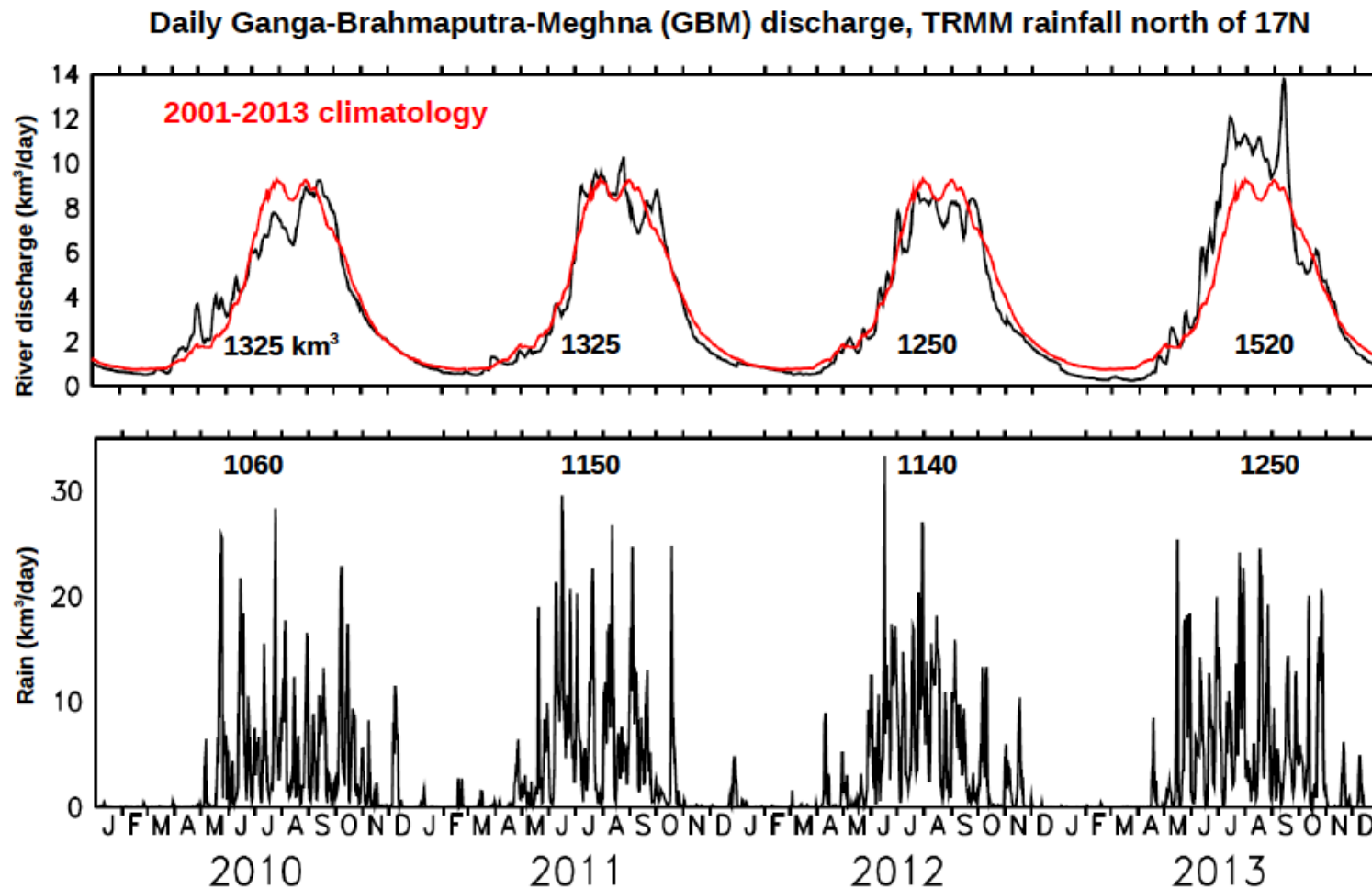
**W. B. Yeats**



Total freshwater input to the Bay of Bengal (P+R-E) 4200 km<sup>3</sup> or 1.6 m





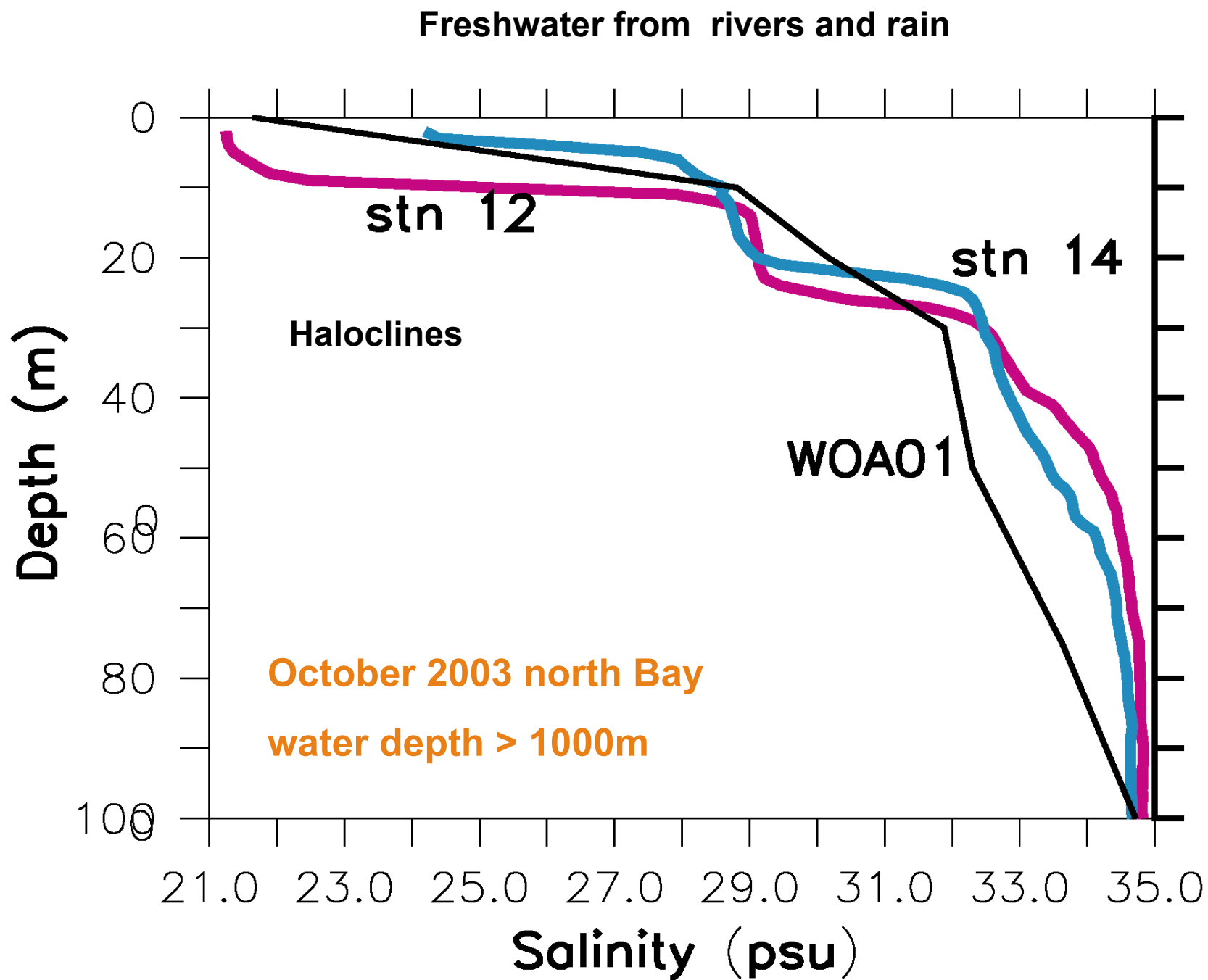


Discharge data courtesy: S. Bala, Fabrice Papa

Institute of Water and Flood Management, Bangladesh University of Engineering and Technology, Dhaka Papa 2012

**Nearly 70 % of fresh water input in June-September**

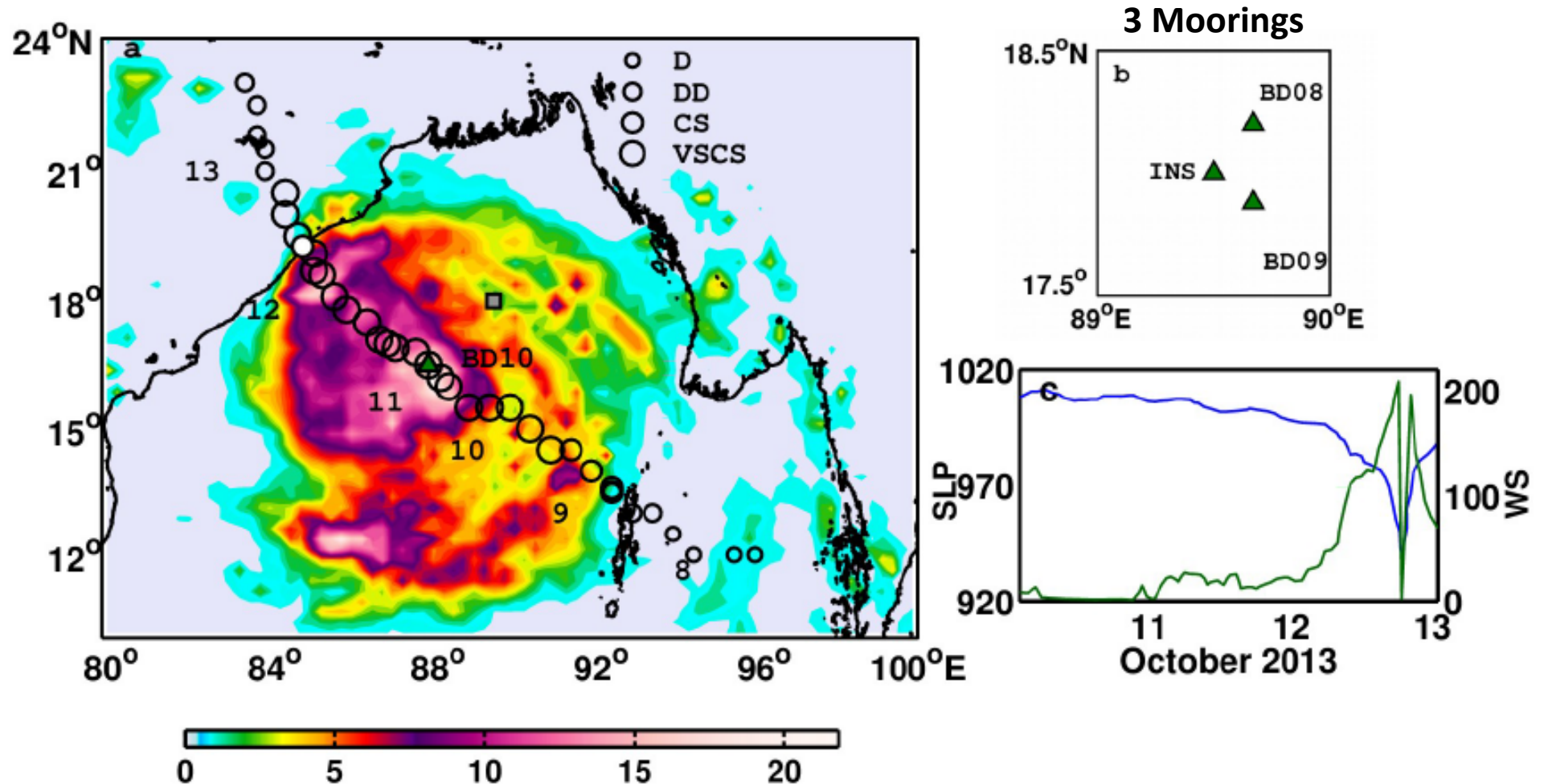






## Tropical cyclone Phailin

8-13 October 2013



940 hPa, 215 km per hour sustained wind speed Gopalpur, Odisha

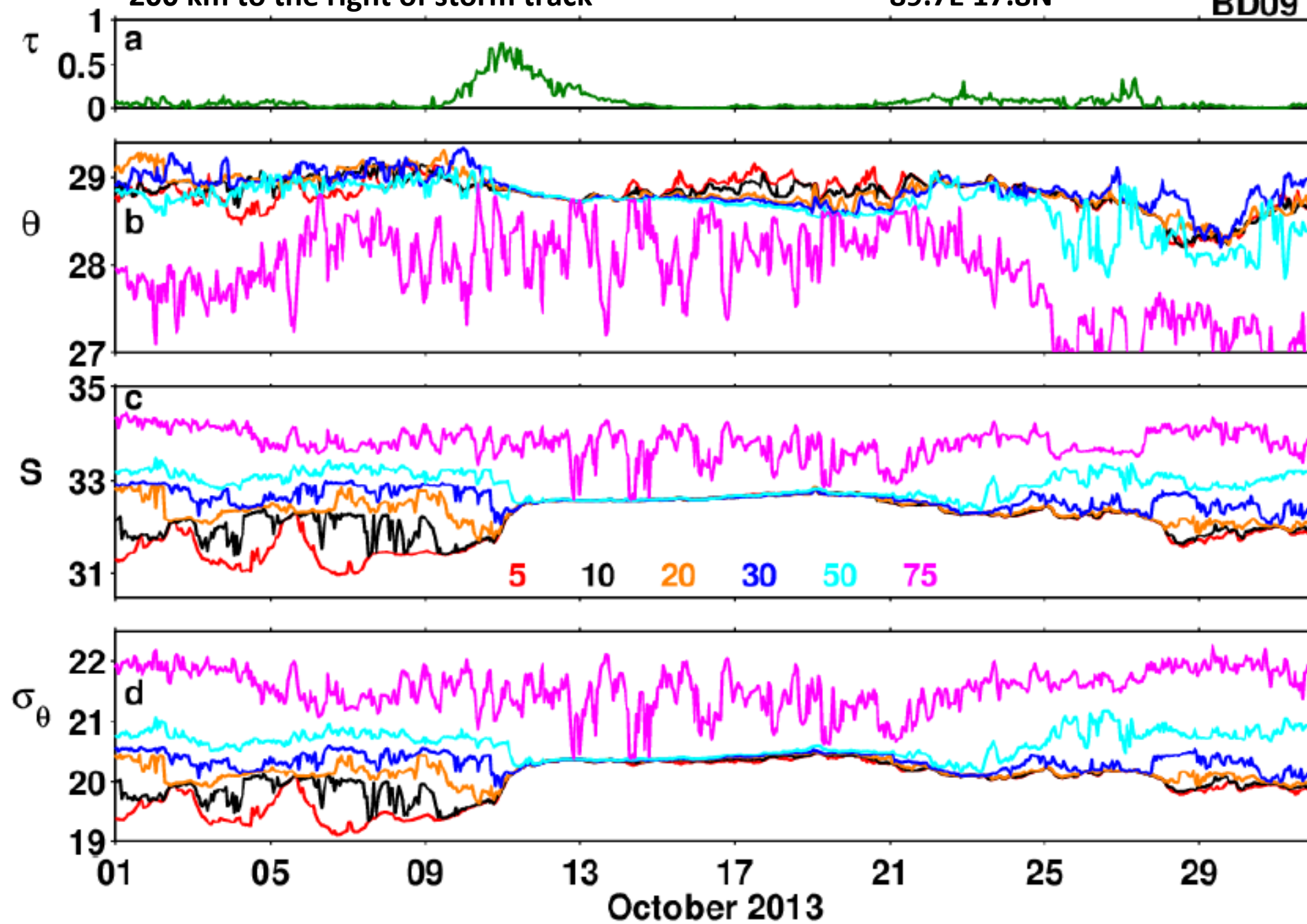
*India Met Department Report 2013*



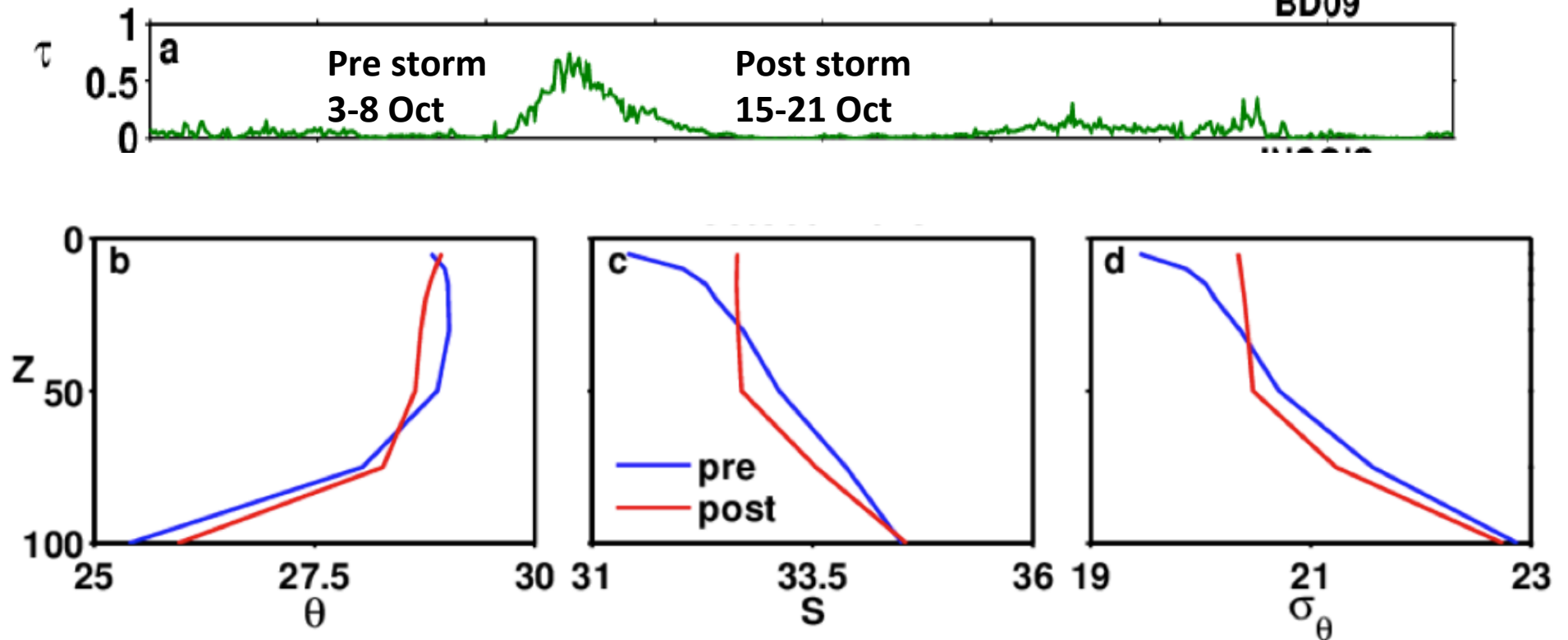
200 km to the right of storm track

89.7E 17.8N

BD09



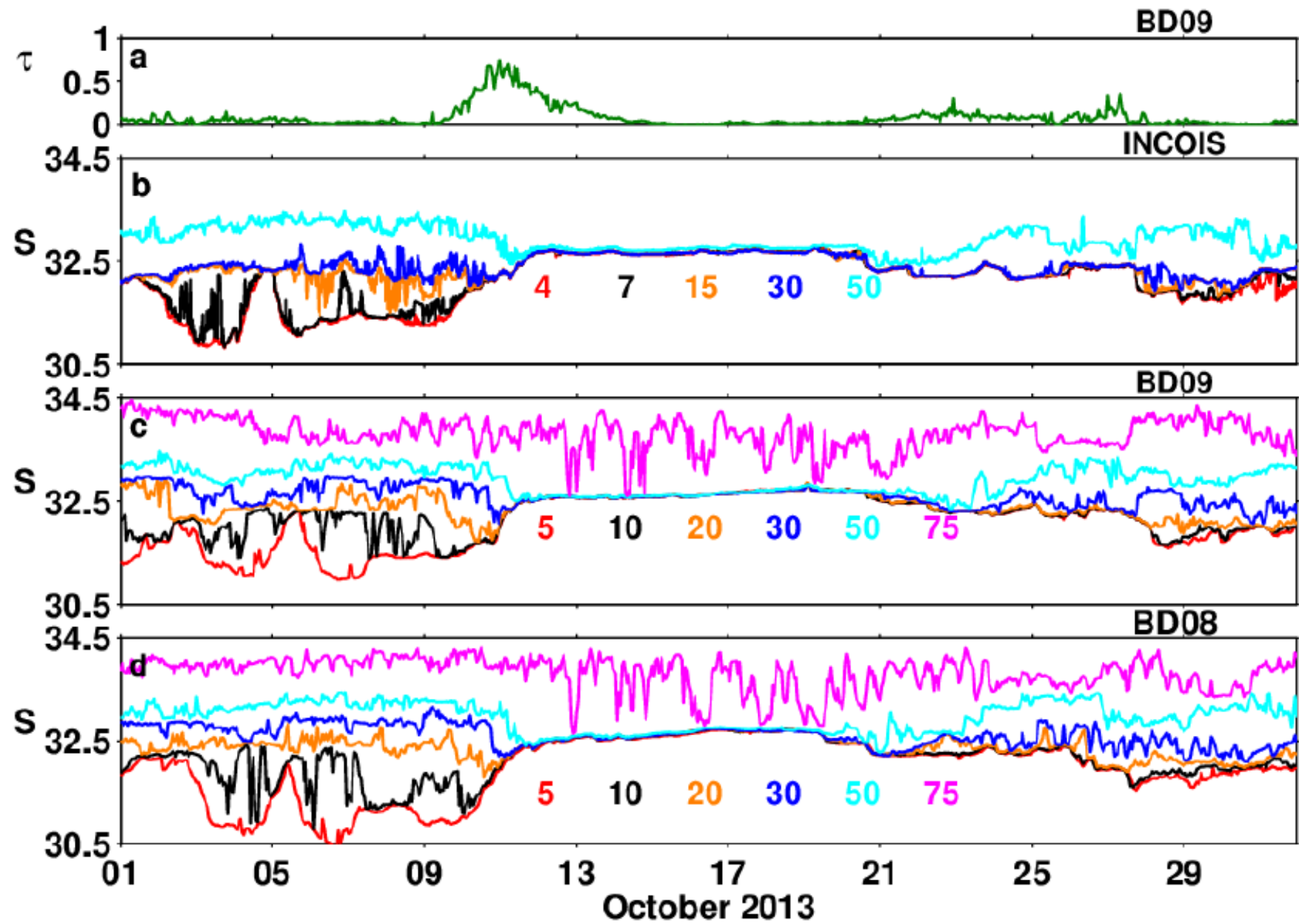
BD09



Thin, fresh surface layer above deep warm layer

Storm-induced vertical mixing: SST increases slightly    SSS increases 1.6 psu

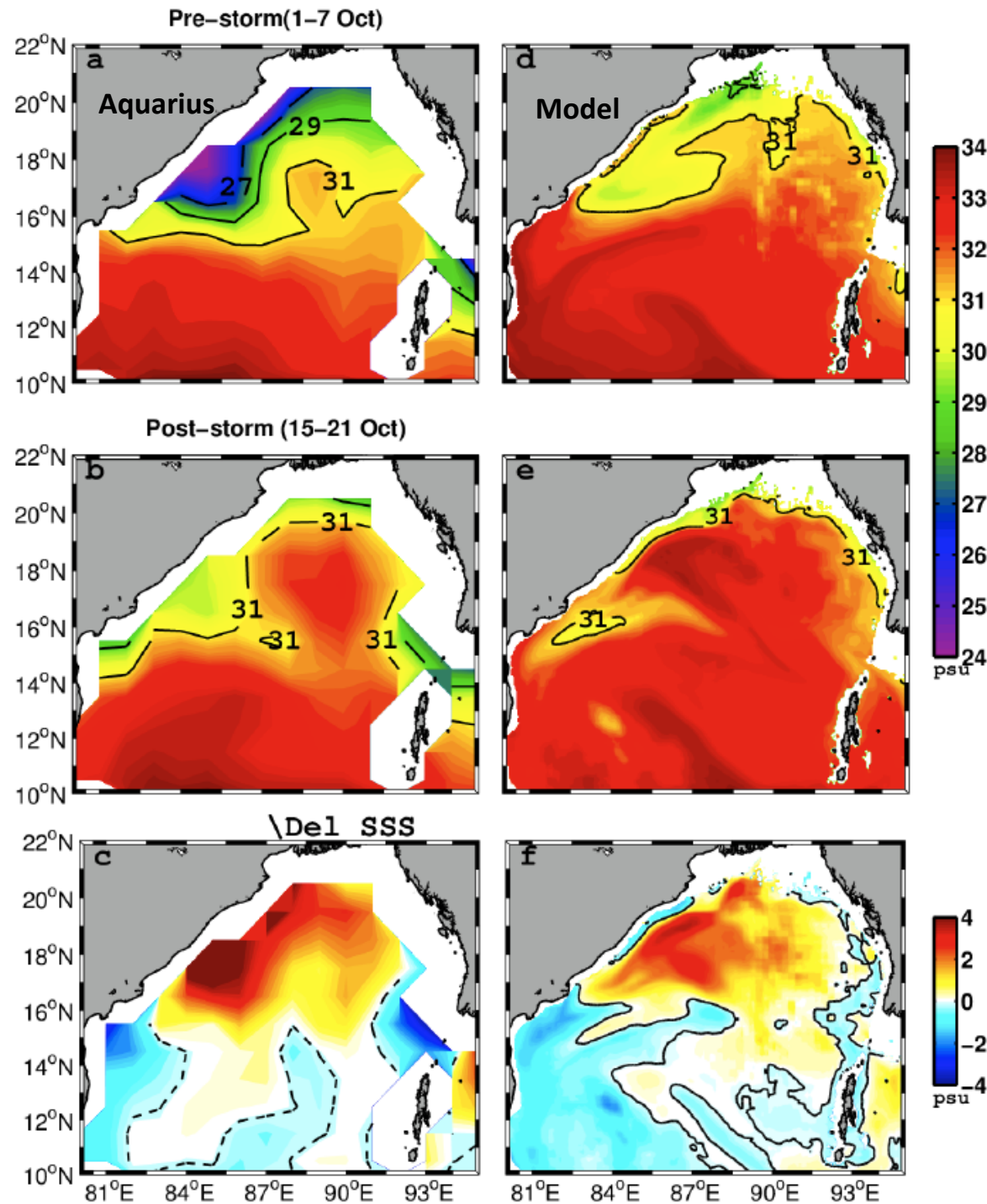




*Sengupta 2008*

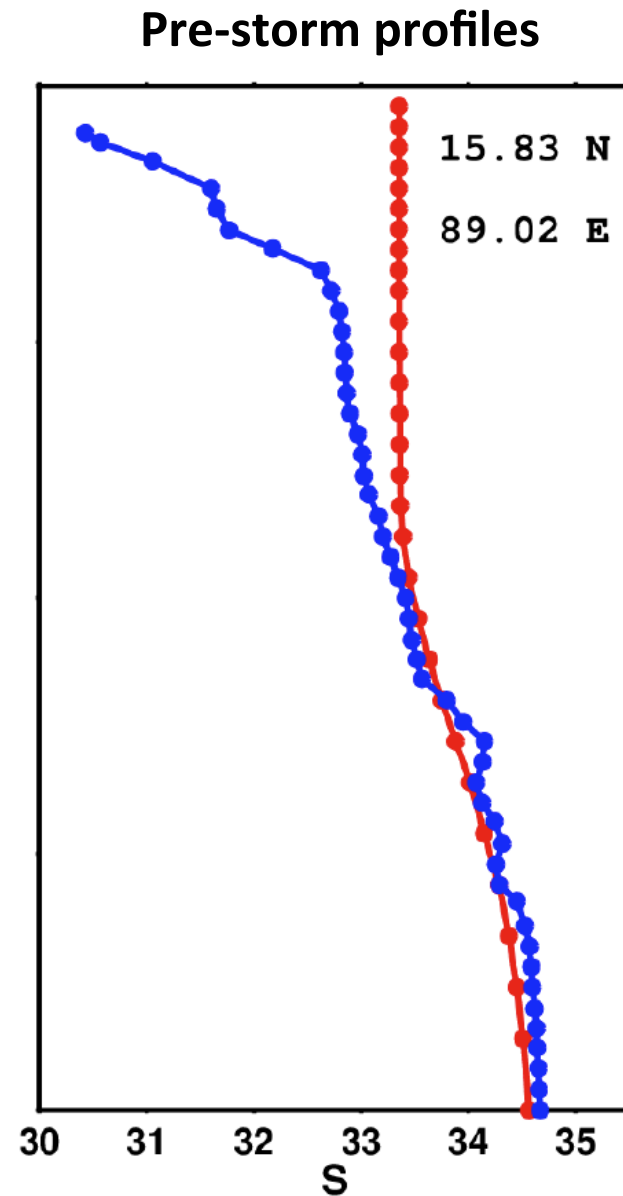
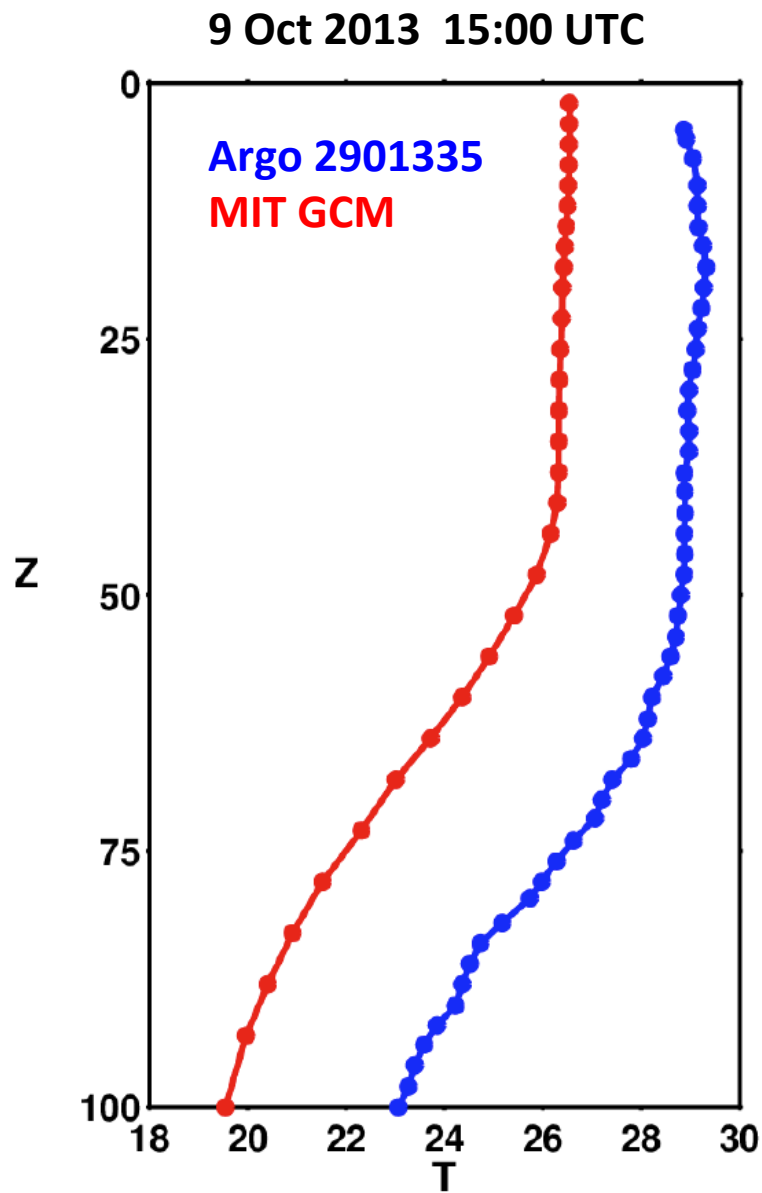
*Singh 2012*

*Balaguru 2012*



**Aquarius satellite: SSS rises  
1.6 psu at mooring location**

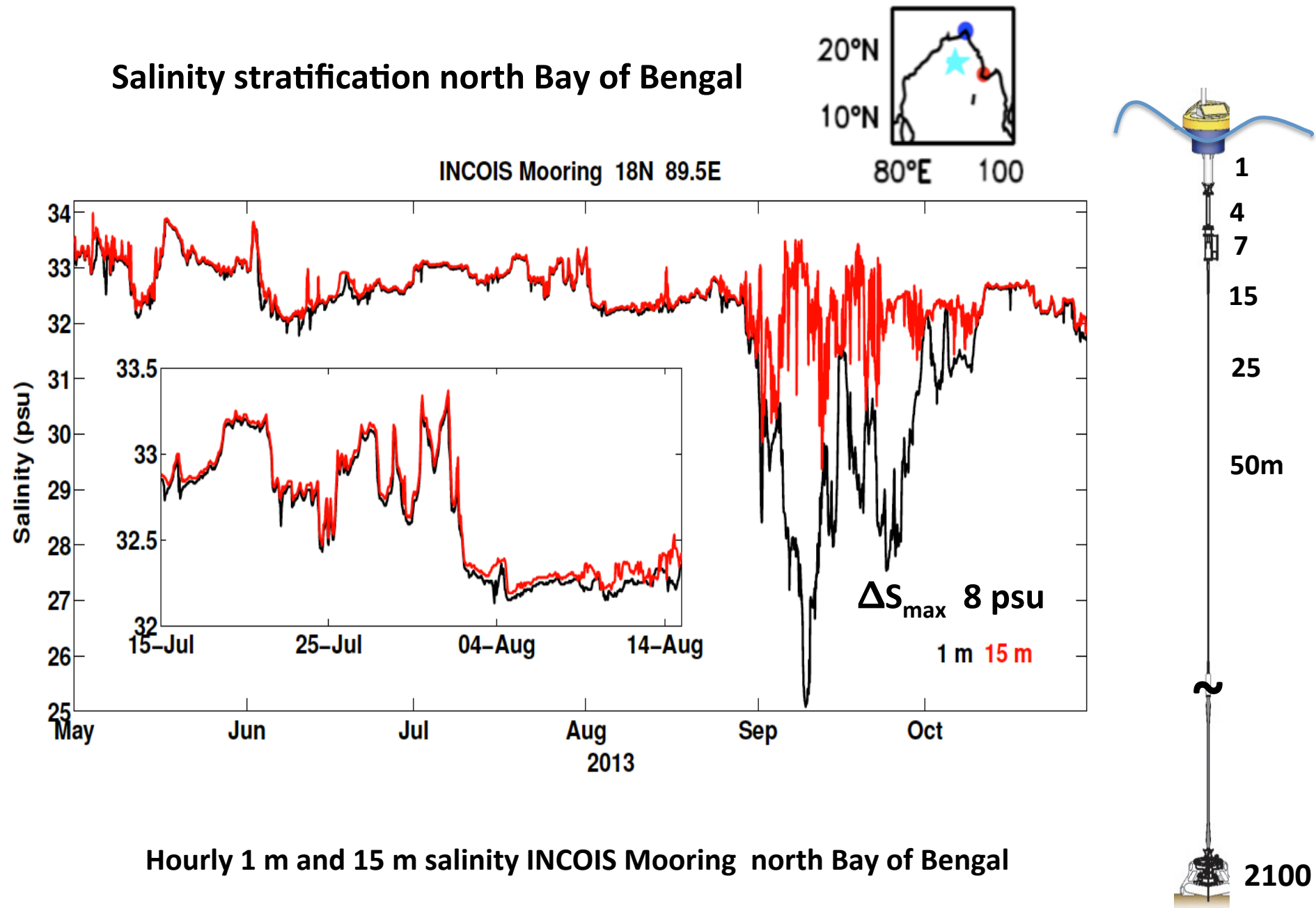
**MIT GCM 5 km, 2 m res.  
Courtesy: *Suneet Dwivedi*  
U. Allahabad**



Near-surface salinity stratification is too weak, and  
upper mixed layer is too deep in most ocean models



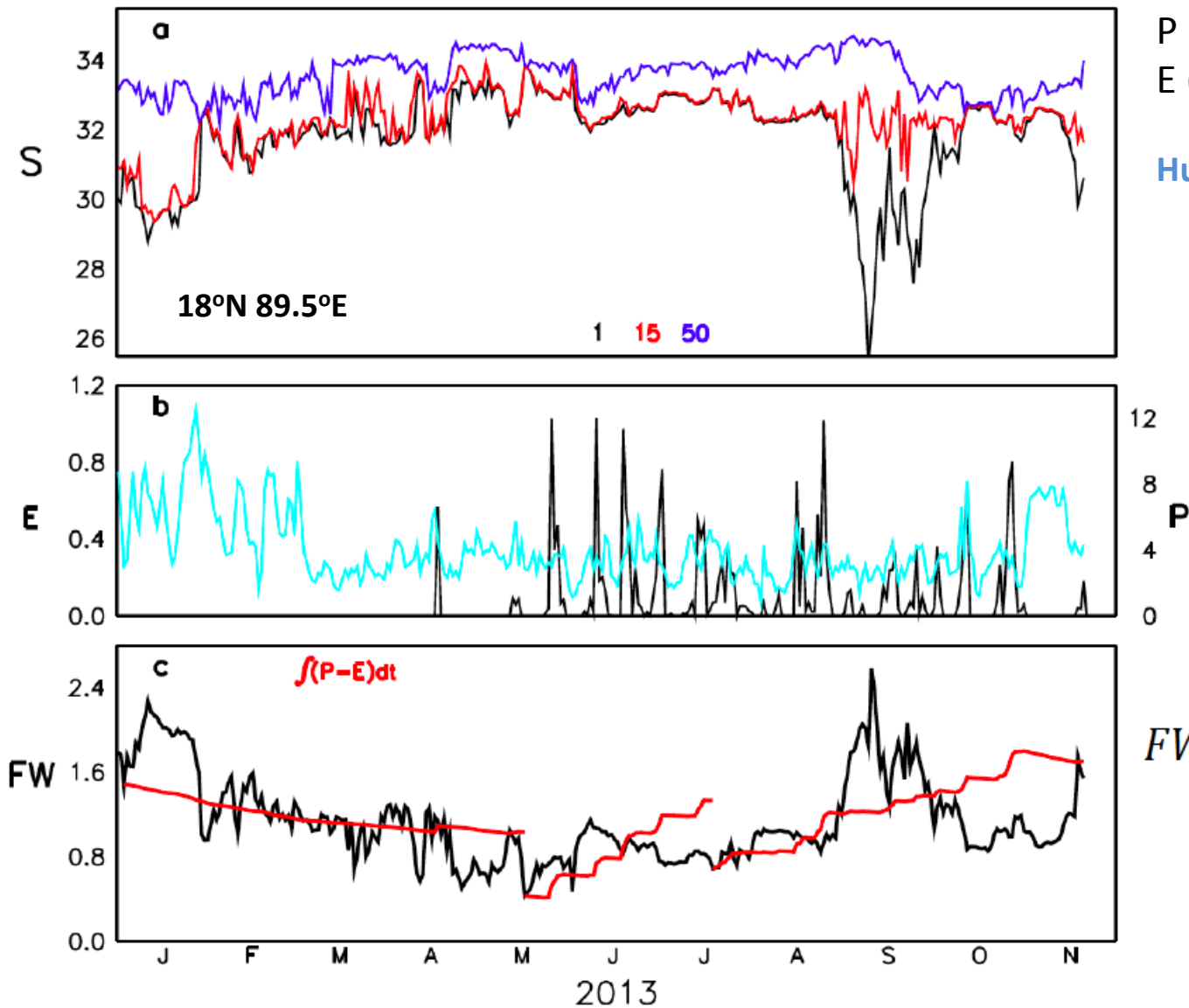
## Salinity stratification north Bay of Bengal



Hourly 1 m and 15 m salinity INCOIS Mooring north Bay of Bengal

*What sustains shallow stratification in the face of mixing?*

## Fresh water balance in the upper 15 m from mooring



P (cm/day) TRMM  
E (cm/day) TropFlux

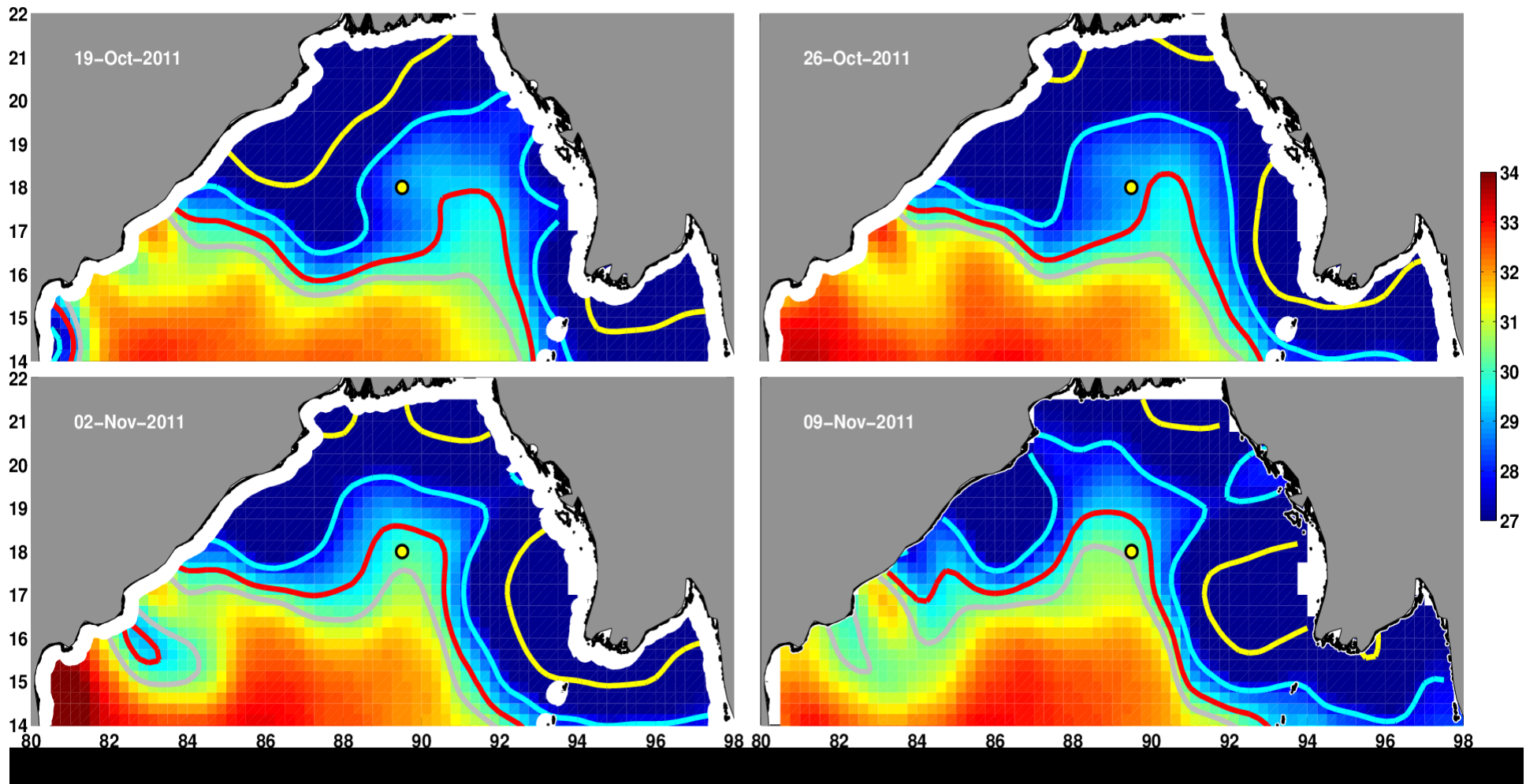
Huffman 2007 Kumar 2012

$$FW = \int_{-15}^0 \left[ 1 - \frac{S(z)}{S_{ref}} \right] dz$$

Pulses of river water at the mooring in summer and winter  
Persistence of river water for three seasons

Sengupta et al., 2016

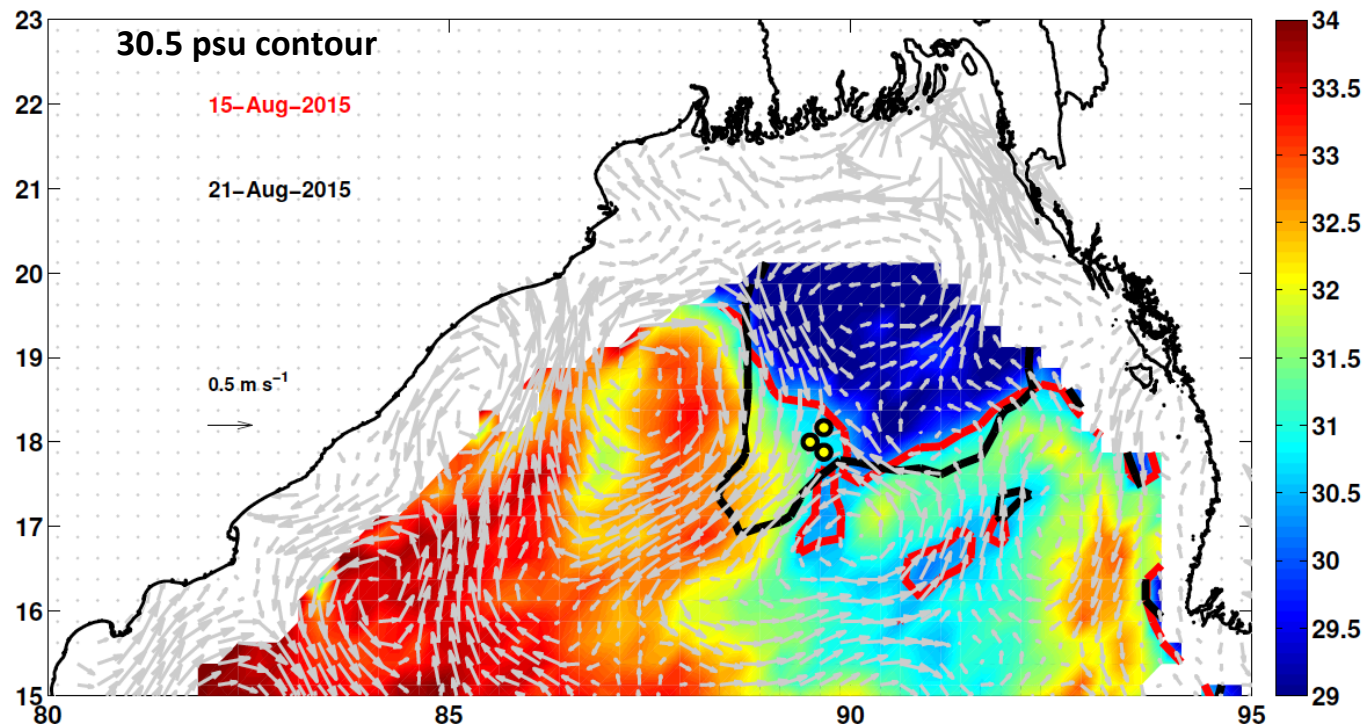
## Aquarius sea surface salinity



30.3 29.3 27.5 24.5 psu contours

Irrawady water from previous year's (2010) summer monsoon discharge





## 300-400 km Eddies

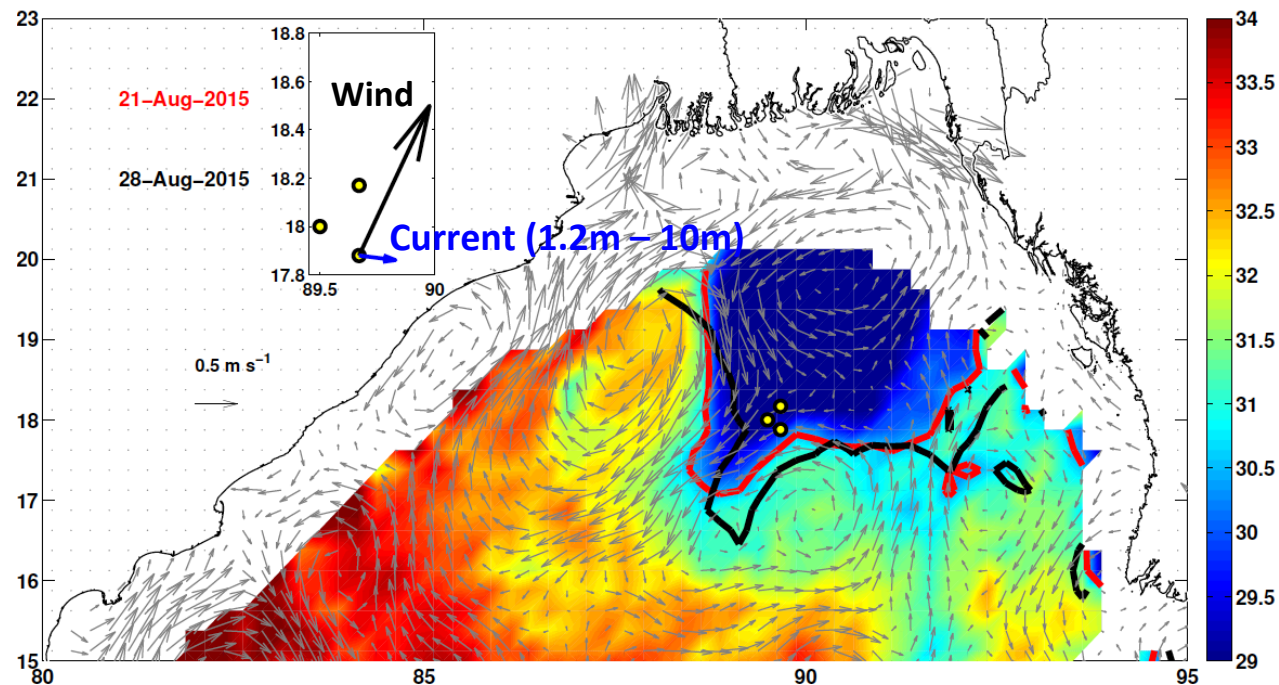
Wind stress 15 Aug 2015  
0.07 N/m<sup>2</sup>

200 km/week due to anti-cyclonic eddy (0.35 m/s)

Wind stress 21 Aug 2015  
0.14 N/m<sup>2</sup>

100 km/week shallow  
Ekman current (0.18 m/s)

## Shallow Ekman Flow

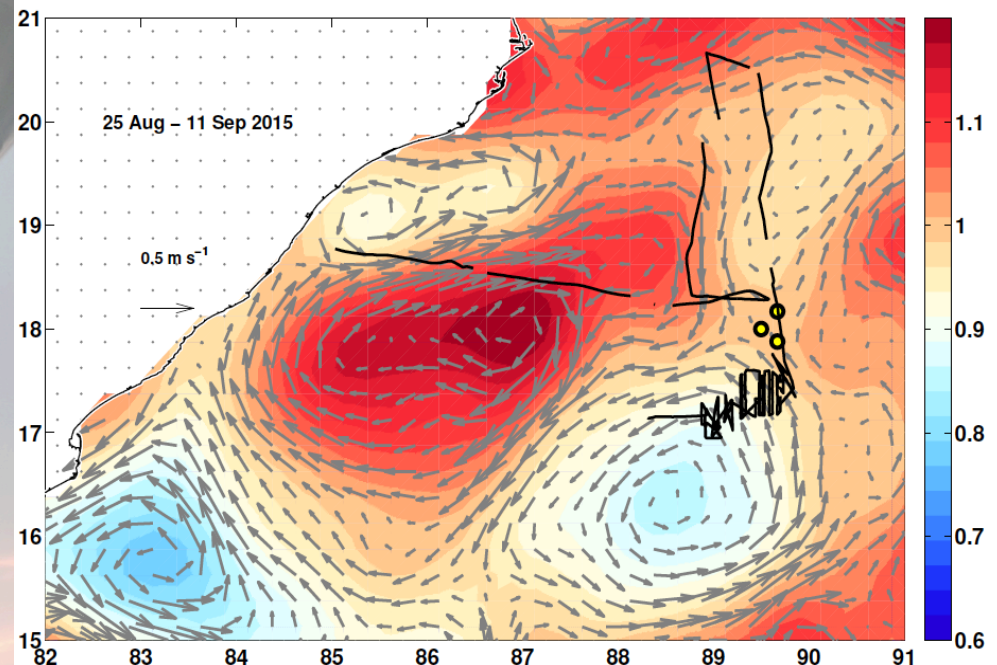
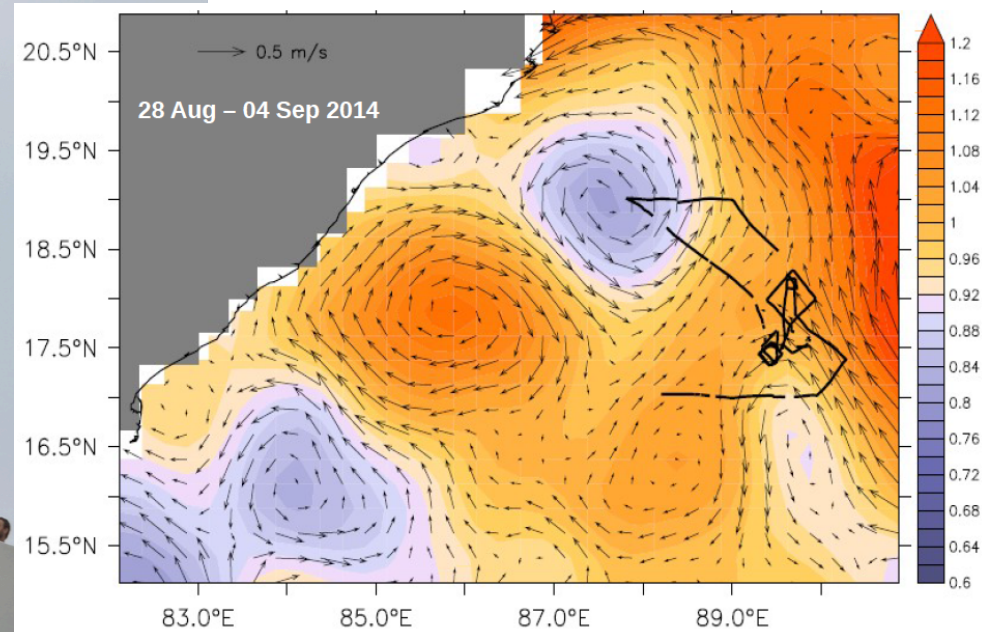


# Two summer monsoon cruises August-September 2014 and 2015

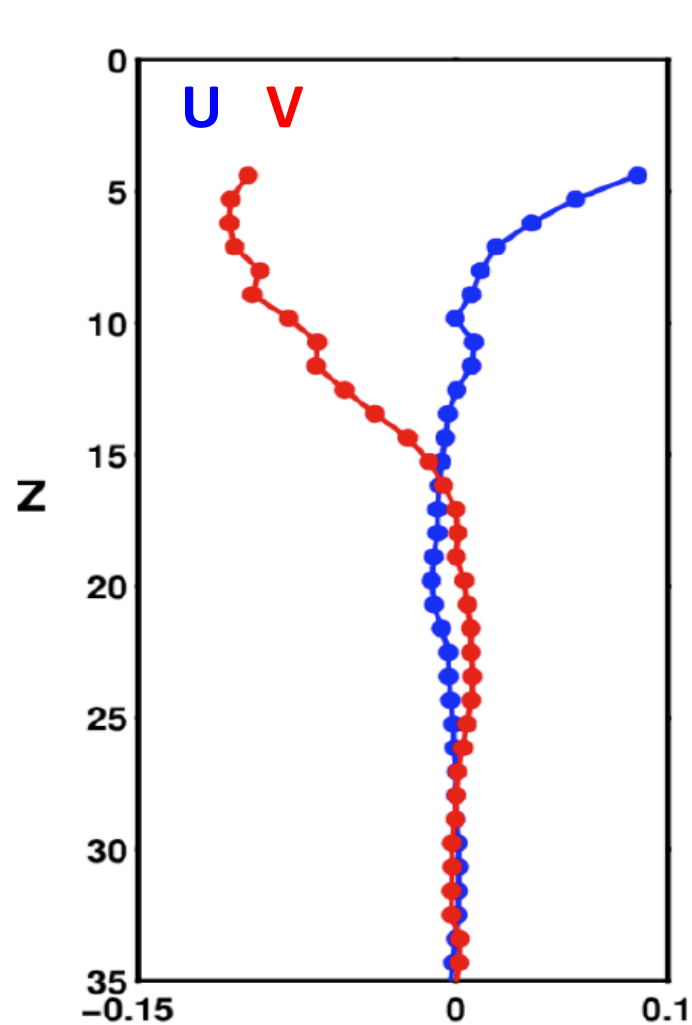
सागर निधि  
SAGAR NIDHI

High res. upper ocean measurements

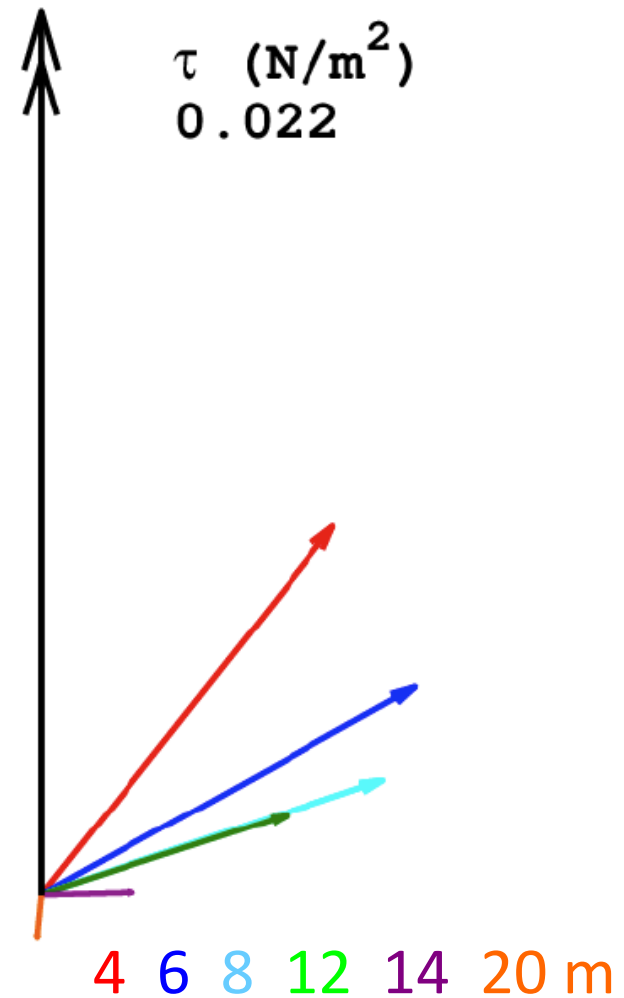
Underway-CTD and ADCP data:  
0.5-1.5 km horizontal res. and 1-2 m  
vertical res.



## Very shallow Ekman spiral



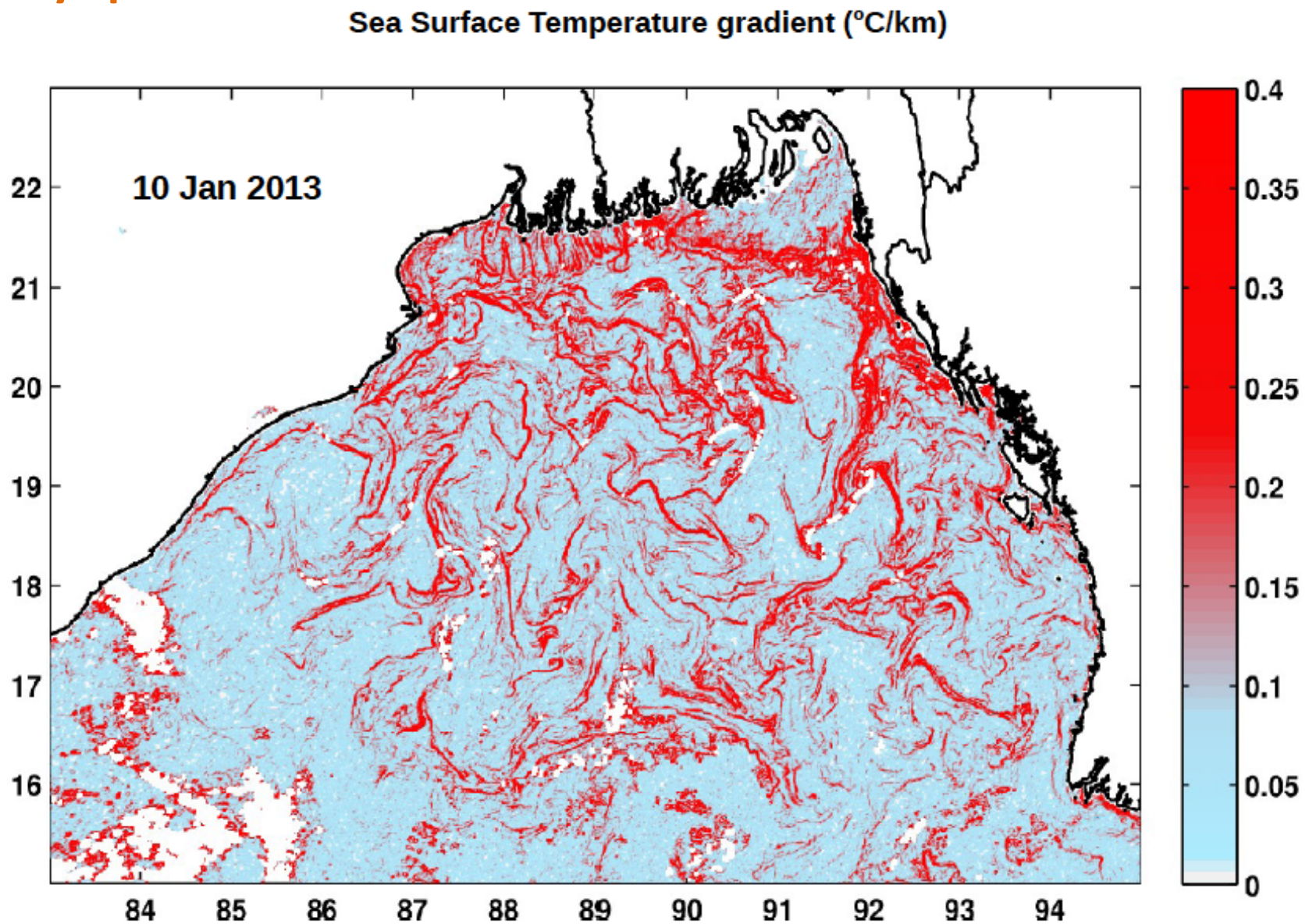
Mean  $U, V$  500 kHz ADCP  
*Sagar Nidhi* 3-11 Sep 2015



Ekman depth about 10m

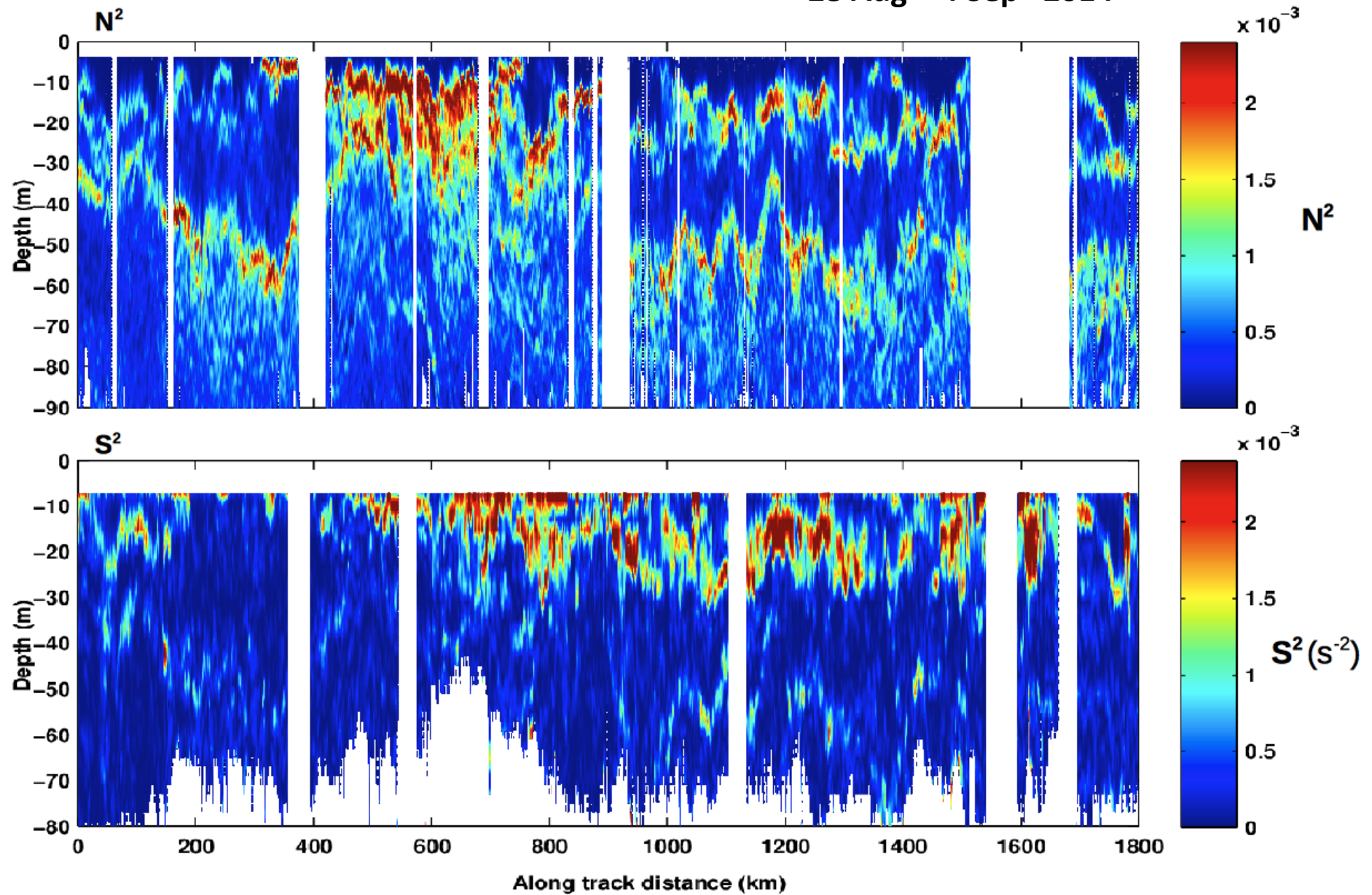


## Many space scales



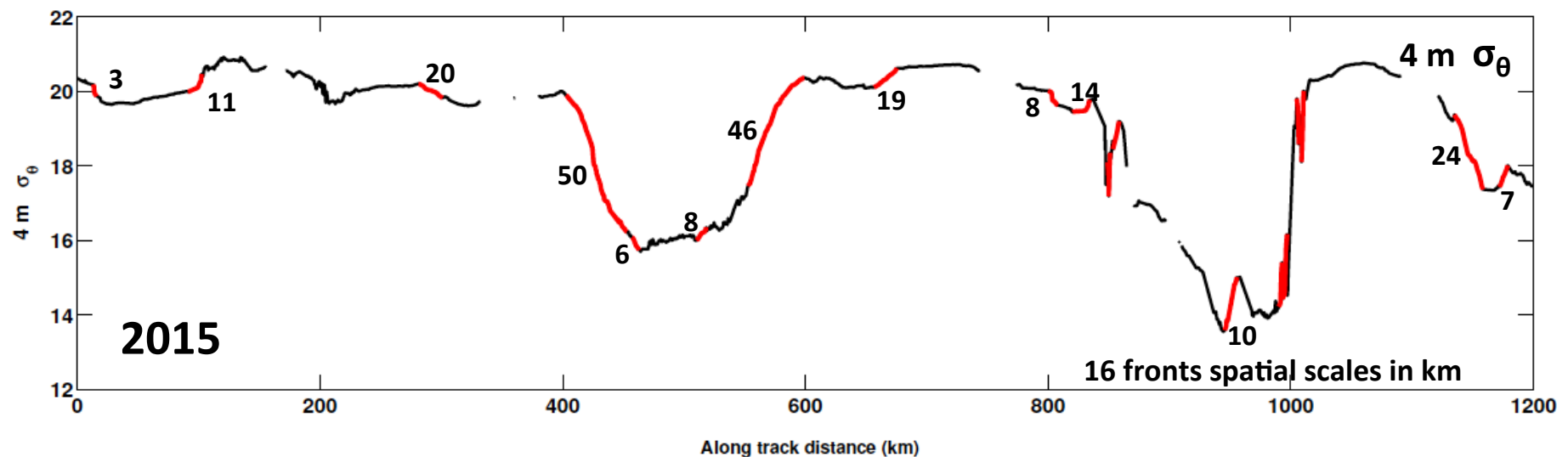
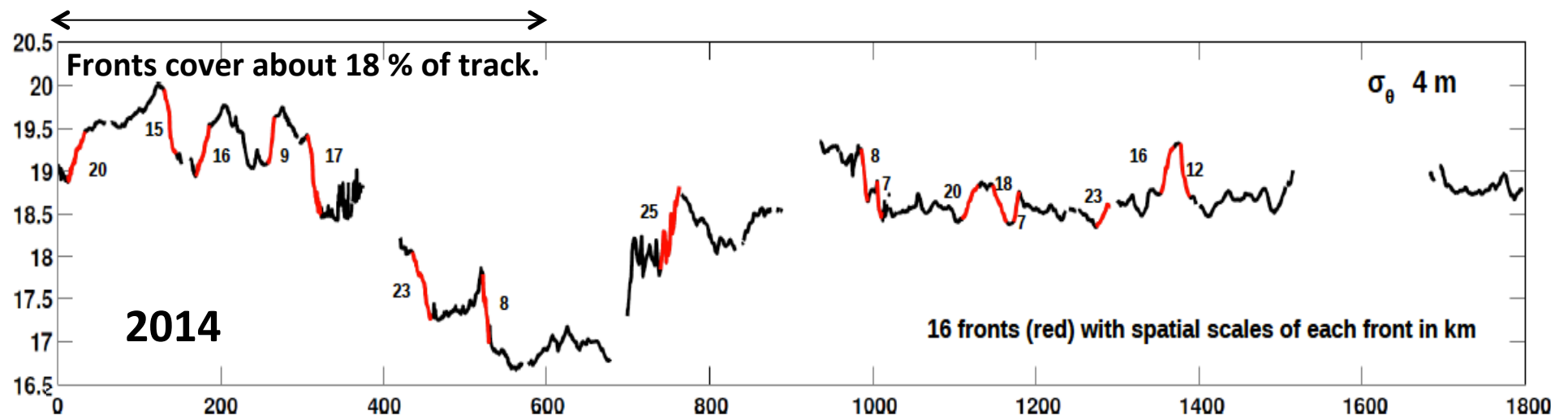
1.1 km resolution NOAA/AVHRR from INCOIS ground station  
Courtesy: Venkat Seshu and Nimit

28 Aug – 4 Sep 2014



$N^2$  (top) and  $S^2 = (du/dz)^2 + (dv/dz)^2$  from ADCP, with 6 m vertical smoothing and 1.4 km horizontal smoothing

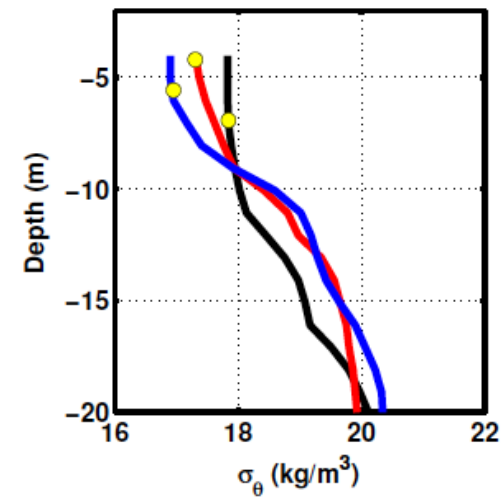
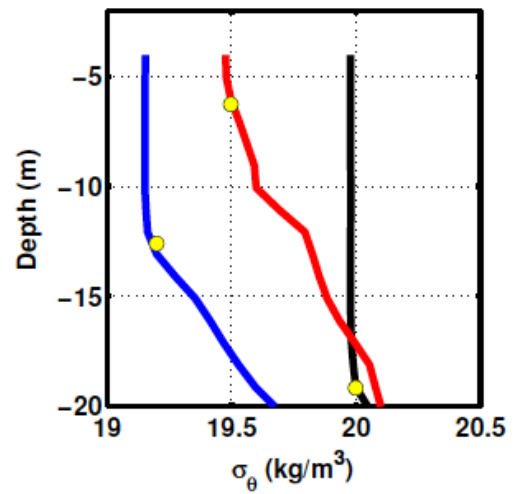
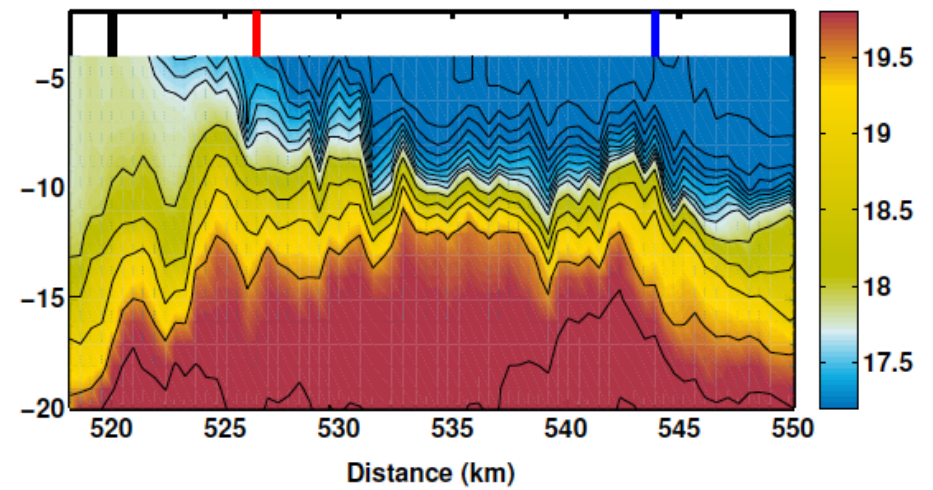
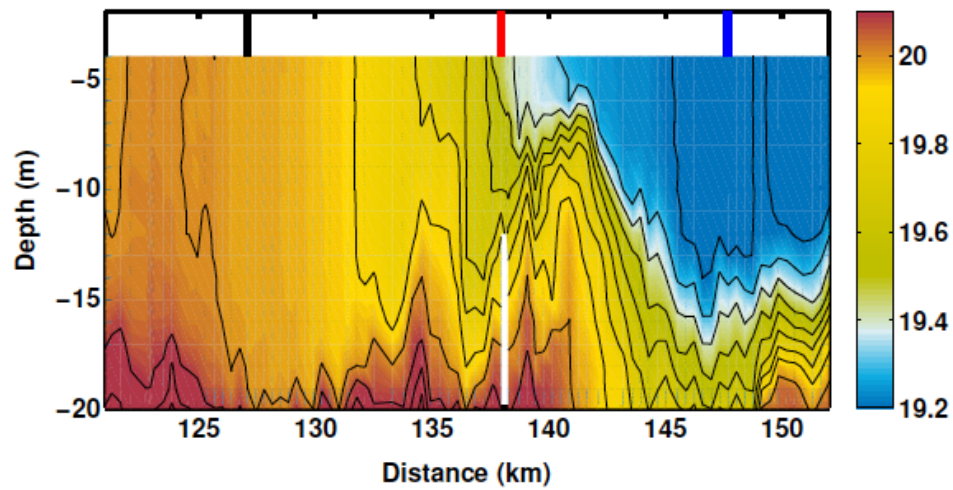




**More than 30 major fronts: Scales 3-50 km, sizes 0.3- 3 kg/m<sup>3</sup>**

*> 10 km data gaps blank*



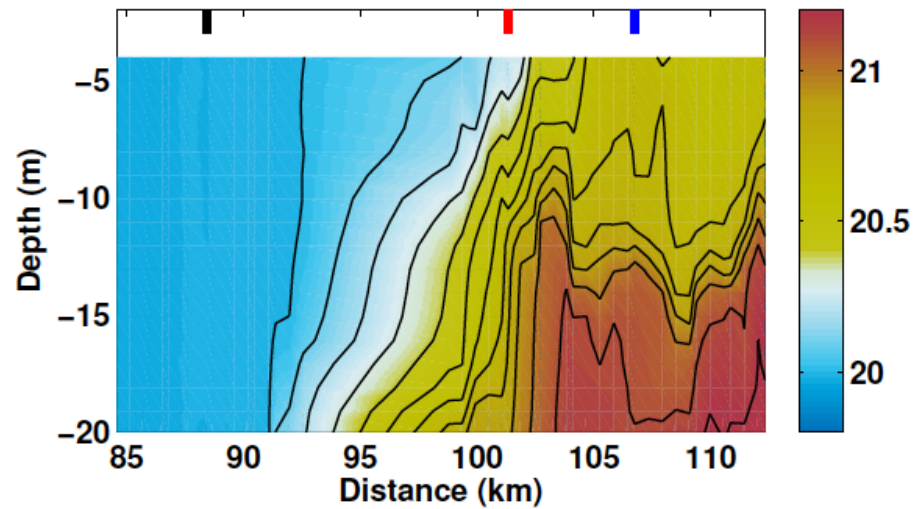


Shallow mixed layers under sub-mesoscale fronts

2014

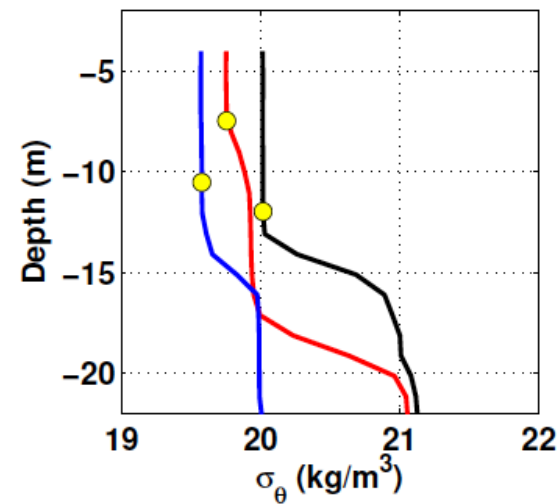
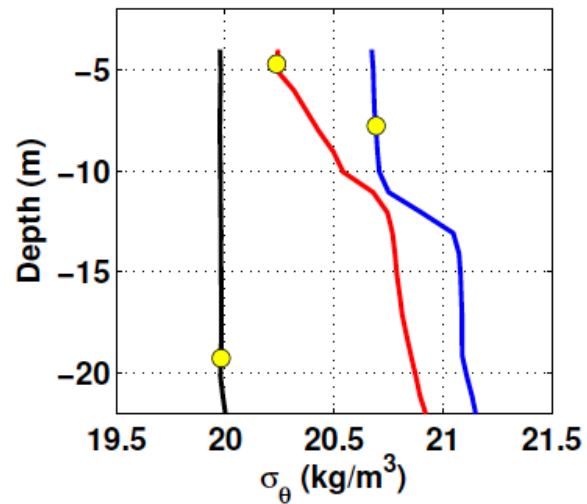
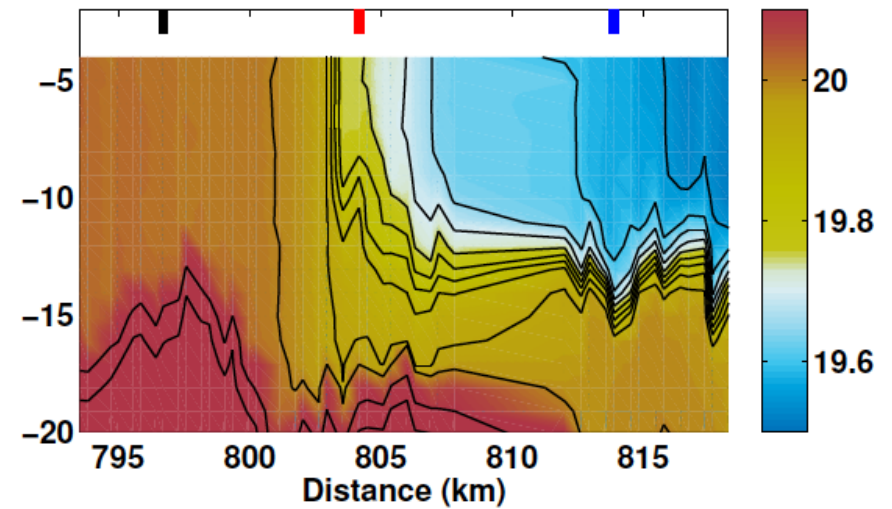
*Timmermans Winsor (2013)* 120 m glider data Arctic

Horizontal density gradient  
0.46 kg/m<sup>3</sup> in 11 km



0.38 kg/m<sup>3</sup> in 7.5 km

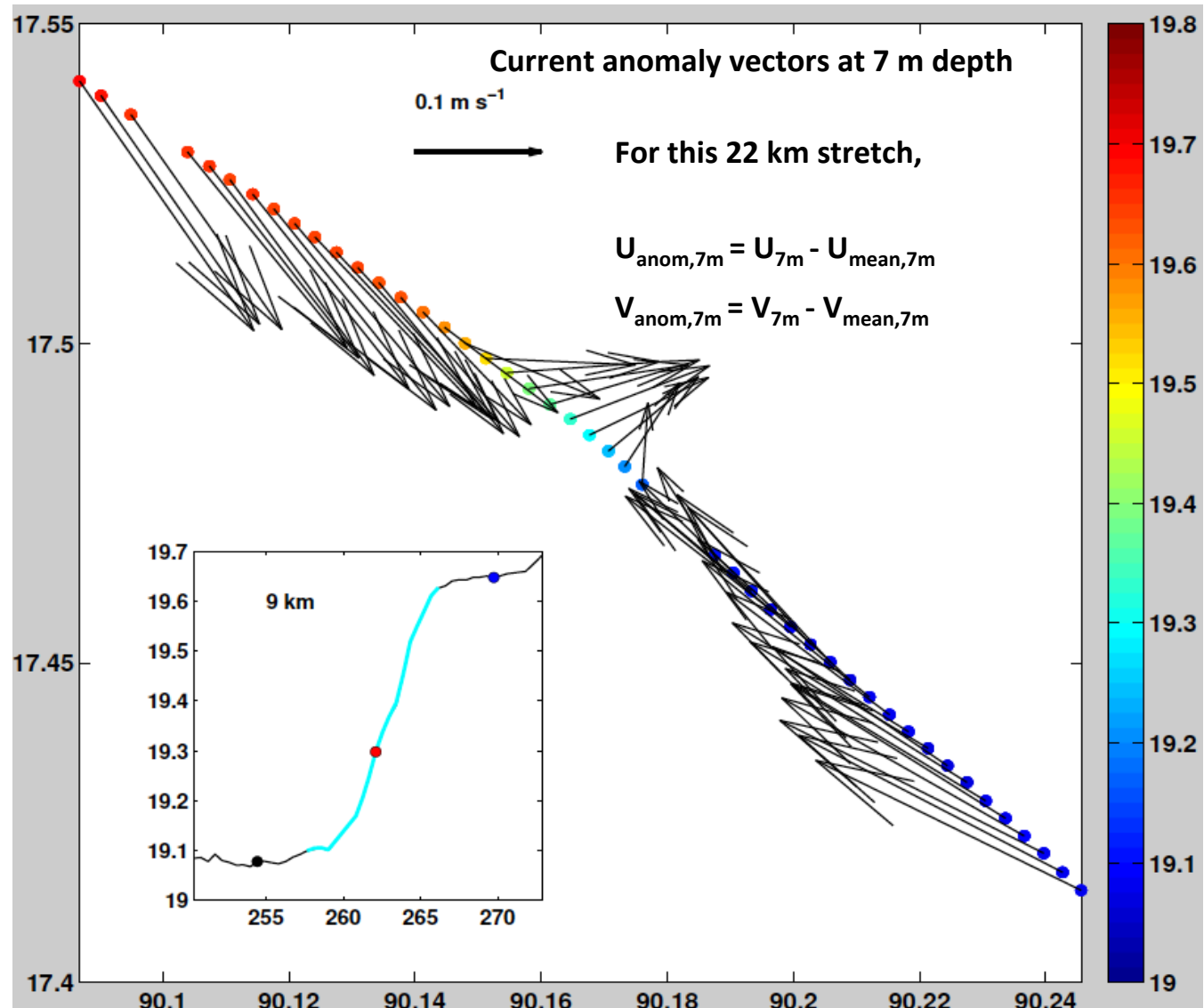
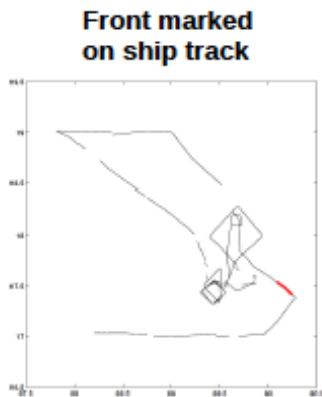
2015



At 17 sub-mesoscale fronts: Shallow mixed layer under the front

## Frontal Jet

$$Ro = \zeta/f \sim 1.5$$

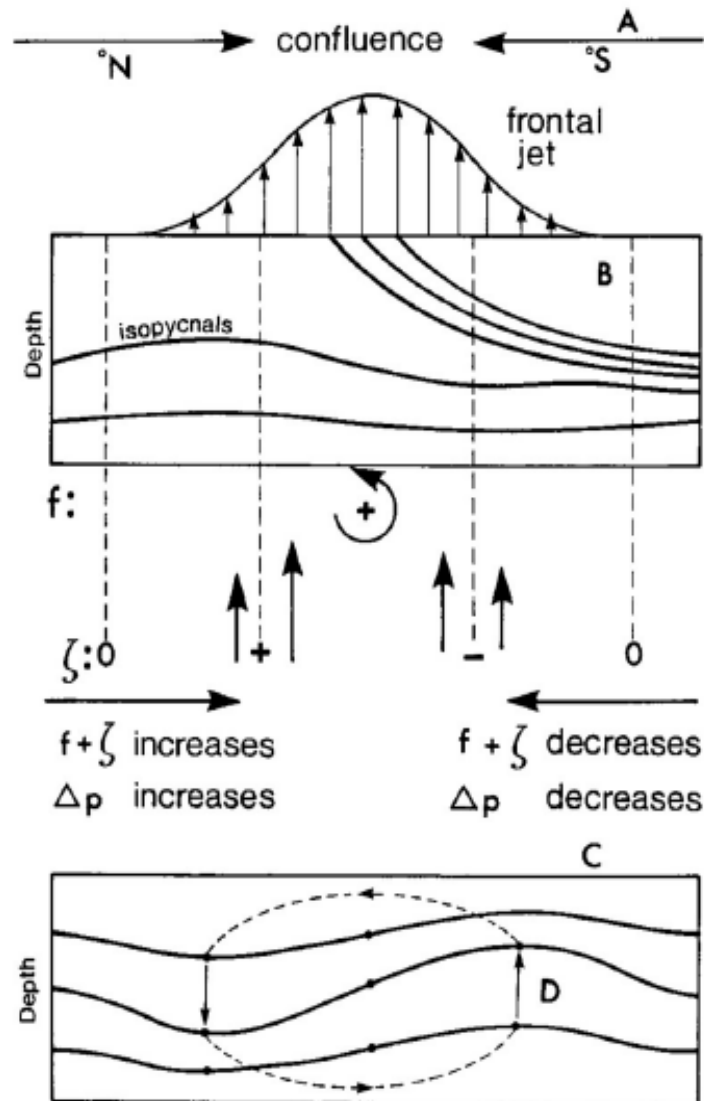


4 m  $\sigma_\theta$  and 7 m velocity anomaly.

Size and lateral scale of front: 0.55 kg/m<sup>3</sup> and 9 km

*Pollard Regier 1991*

## Slumping at sub-mesoscale fronts creates stratification



**Pollard and Regier 1991**

Potential vorticity is conserved

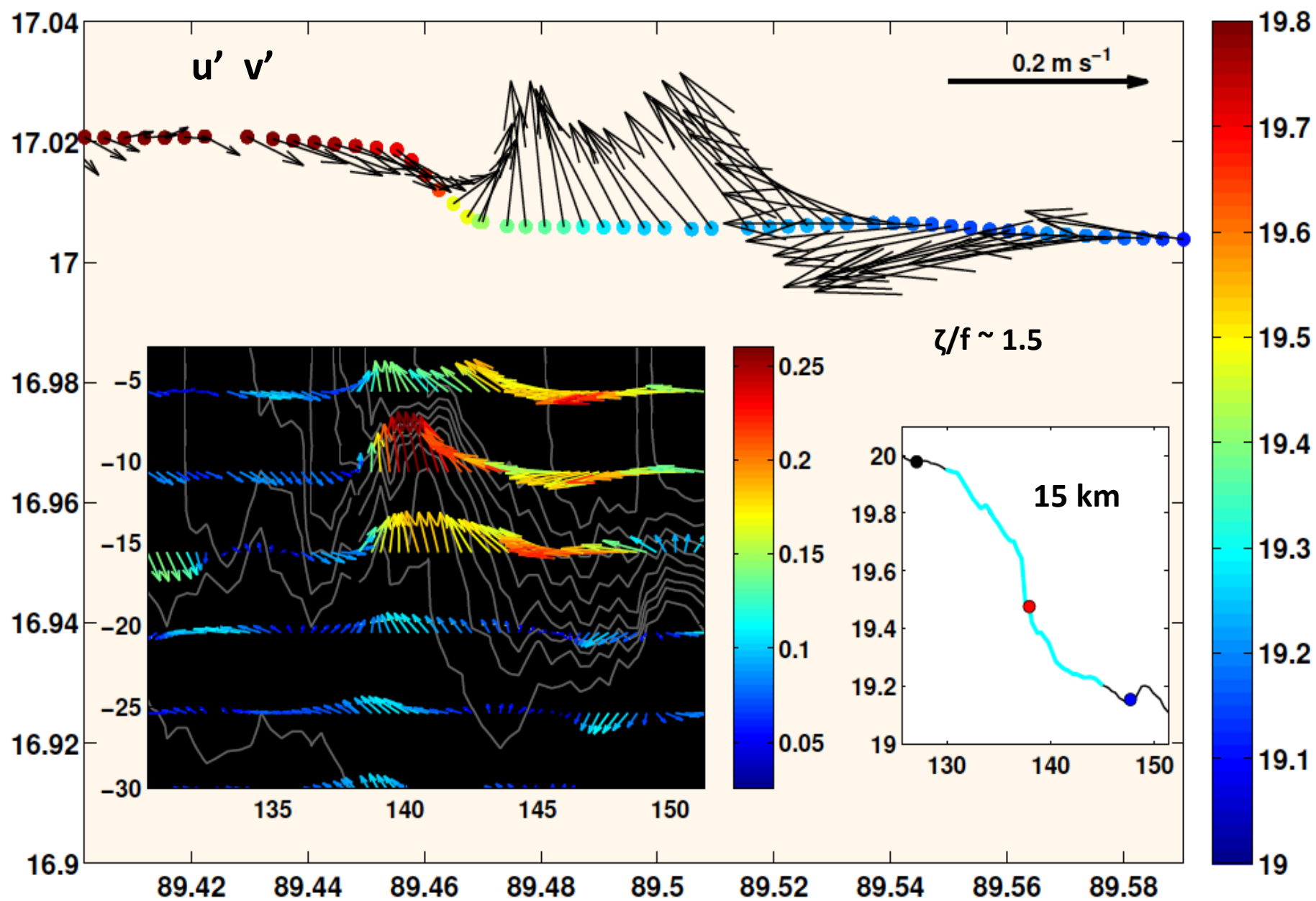
$$D(PV)/Dt = 0$$

$$PV = (f + \zeta) \cdot \text{grad } \rho$$

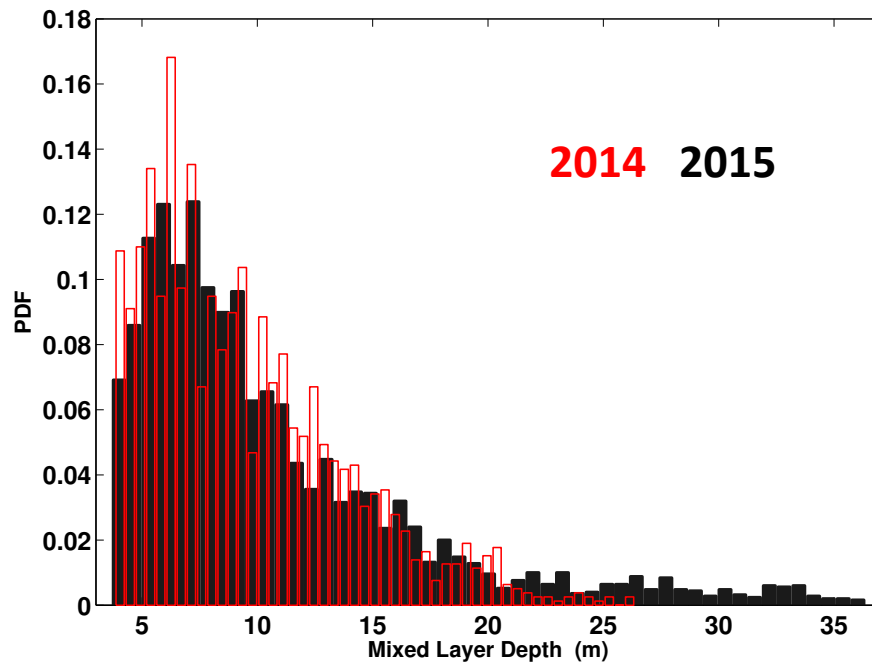
Downwelling on the dense side of the front and  
upwelling on the less dense side of the front

**Tandon and Garrett 1995 Mahadevan and Tandon 2006 LN Thomas 2007**





Frontal jet shallower than 20 m, lateral scale 5-10 km

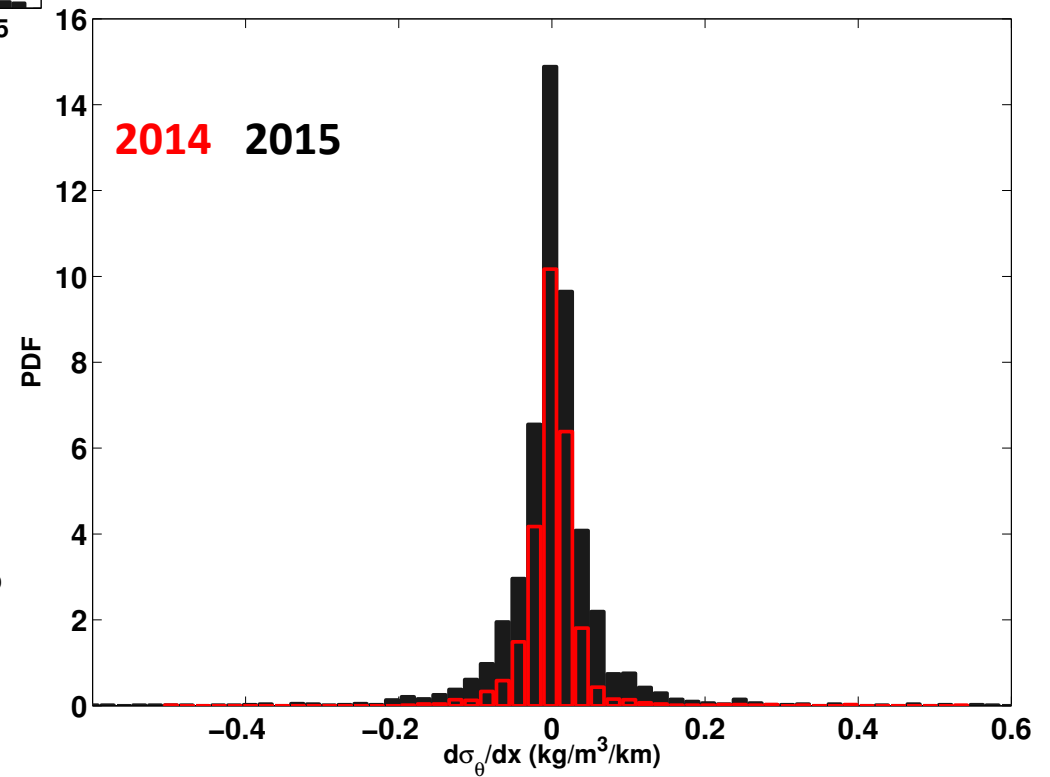


### Mixed Layer Depth

MLD < 10 m about 60 % of the time in both years

### Lateral density gradient

$|\mathrm{d}\sigma_\theta/\mathrm{d}x| > 0.05 \text{ kg/m}^3/\text{km} \sim 10\% \quad 22\%$



## Summary

**Shallow, low-salinity surface layer and deep warm layer** in the north Bay of Bengal is important for regional air-sea interaction. Most **models are inadequate at simulating this thermodynamic structure.**

River water, dispersed by basin-scale circulation,  $O(100)$  km eddies and shallow Ekman flow, **persists for at least three seasons in the north Bay.**

First measurements with  $O(1\text{ km})$  resolution in this basin show near-surface **salinity-dominated “sub-mesoscale” fronts with 1-10 km lateral scales.**

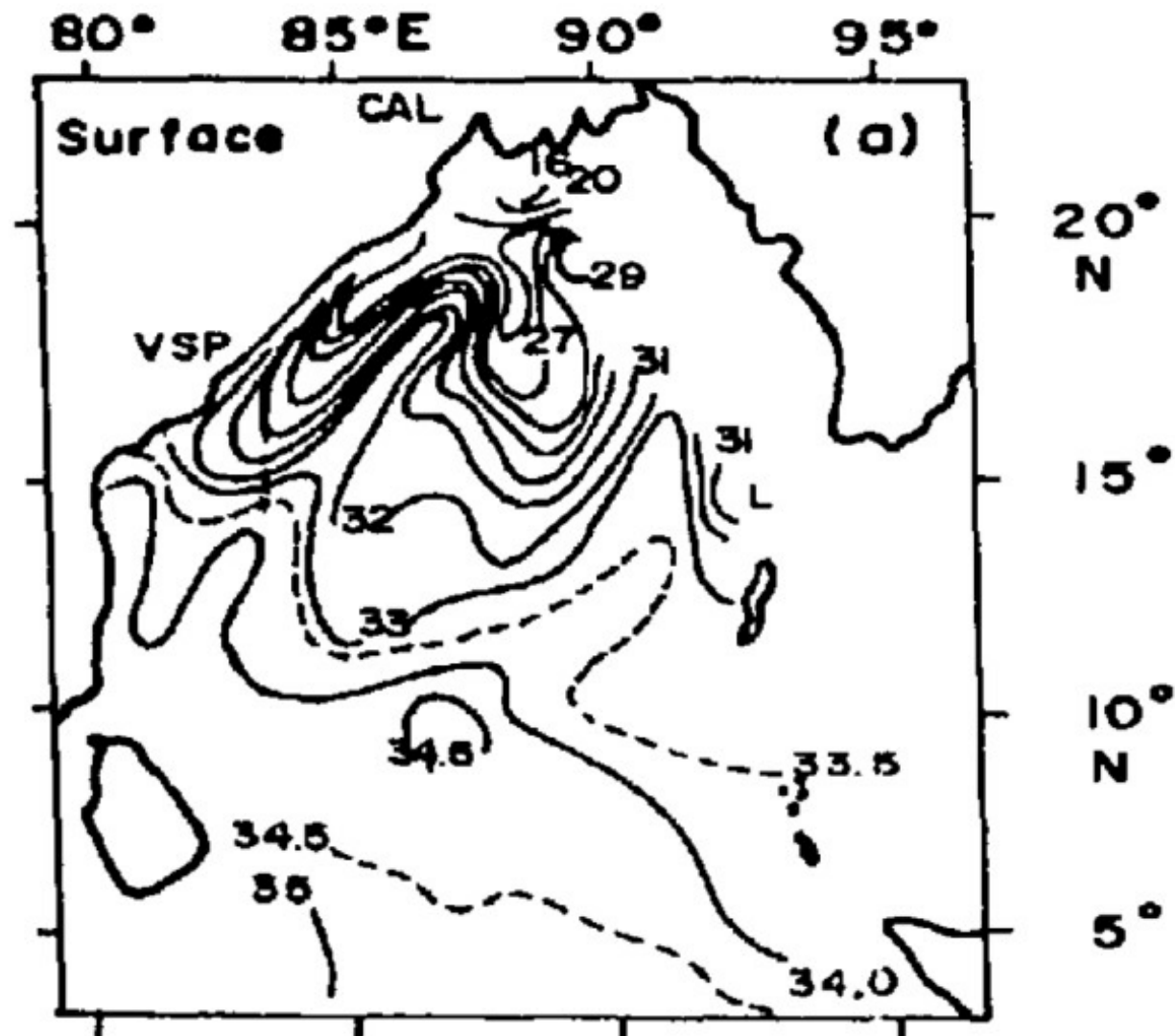
**Mixed layer depth** is often shallower at the 1-10 km sub-mesoscale fronts.

**Slumping of sub-mesoscale fronts may sustain shallow stratification**





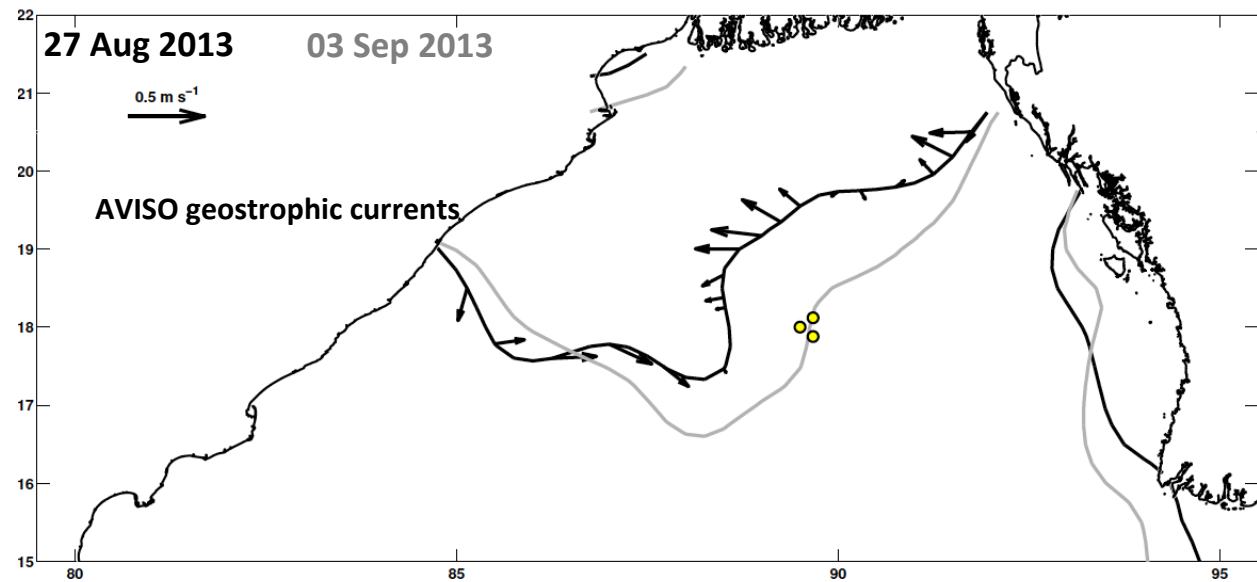
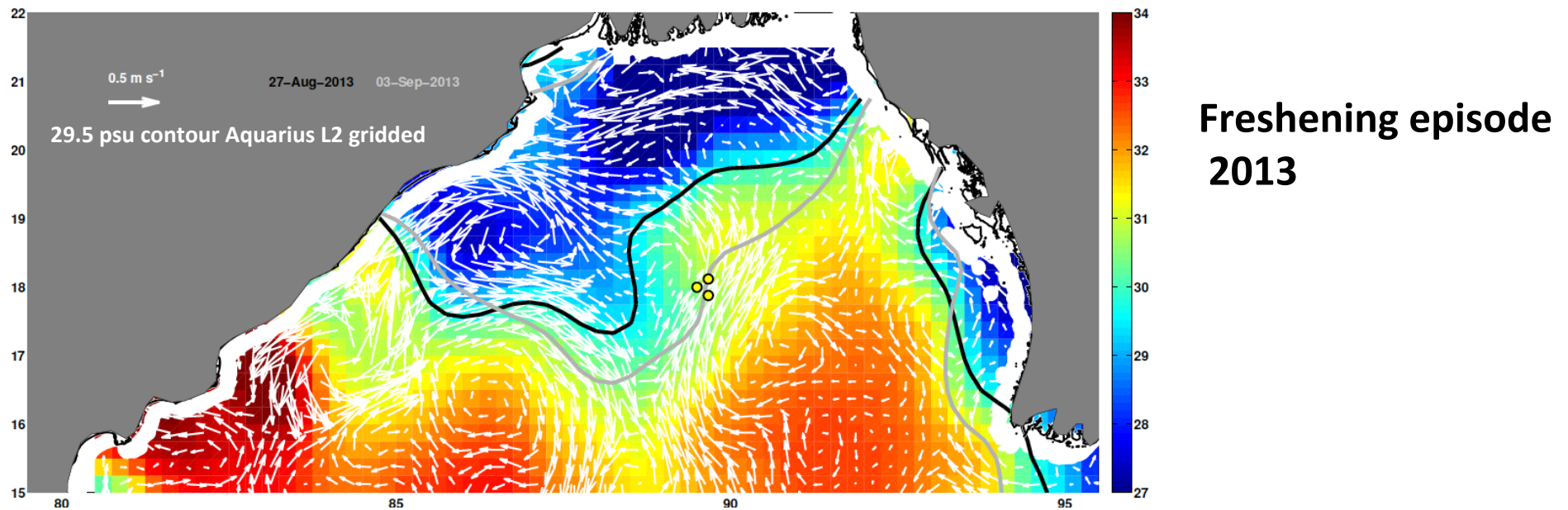
## River plumes, July-September Sea Surface Salinity



Sagar Kanya cruise 1984

*Murty et al., 1992*

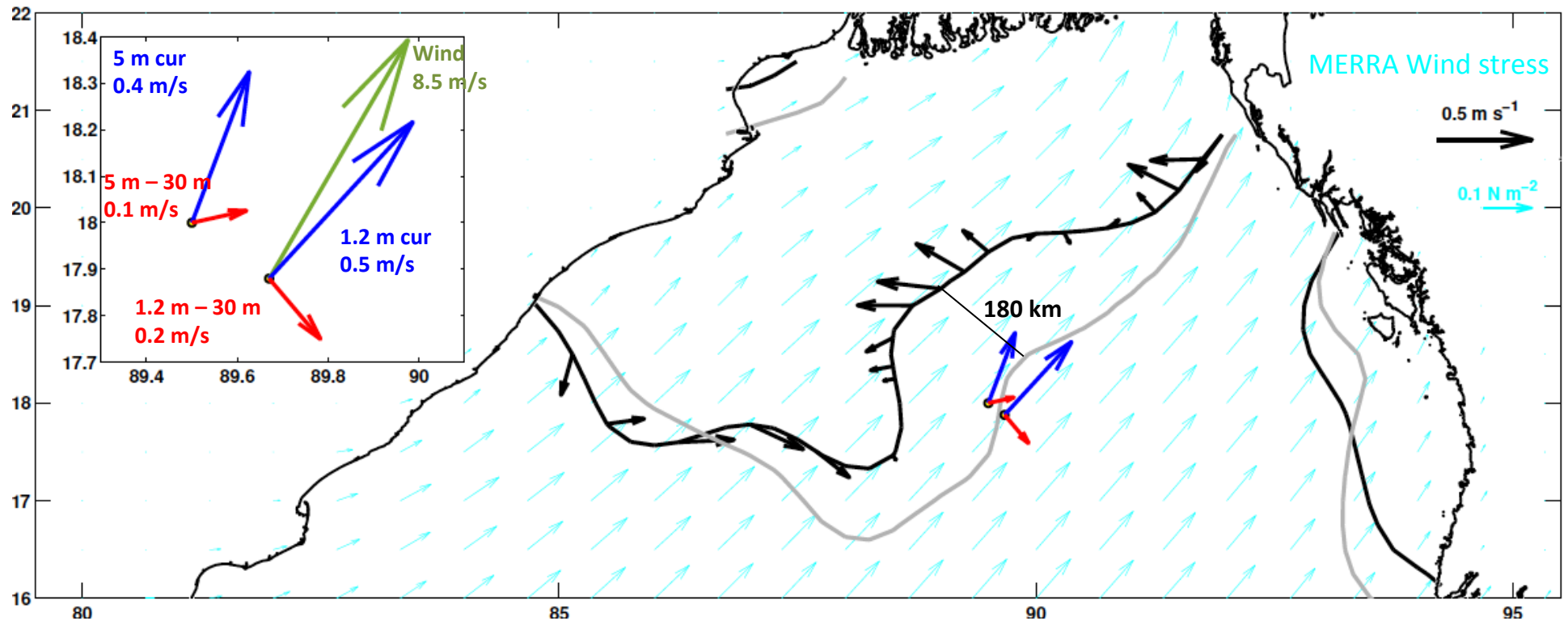
# Mechanisms of freshwater movement: Ekman transport



**Why does the isohaline move to the south and east?**

Contour displaced by velocity component normal to the isohaline.

Ekman velocity at 1.2 m depth is 0.2 m/s. Displacement in a week is 120 km

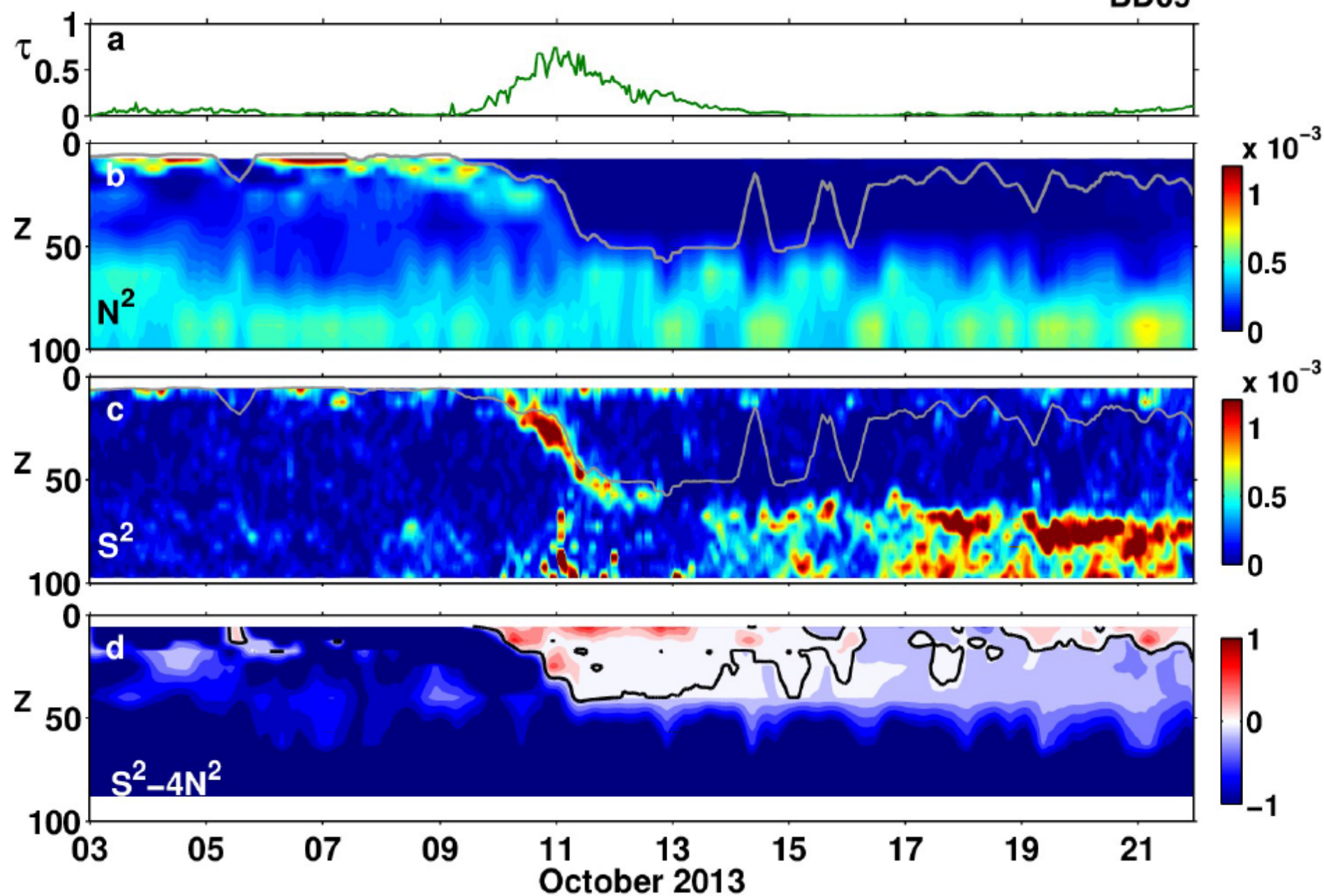


$$\text{Ekman Transport} = \tau / \rho f = 2.2 \text{ m}^2/\text{s}$$

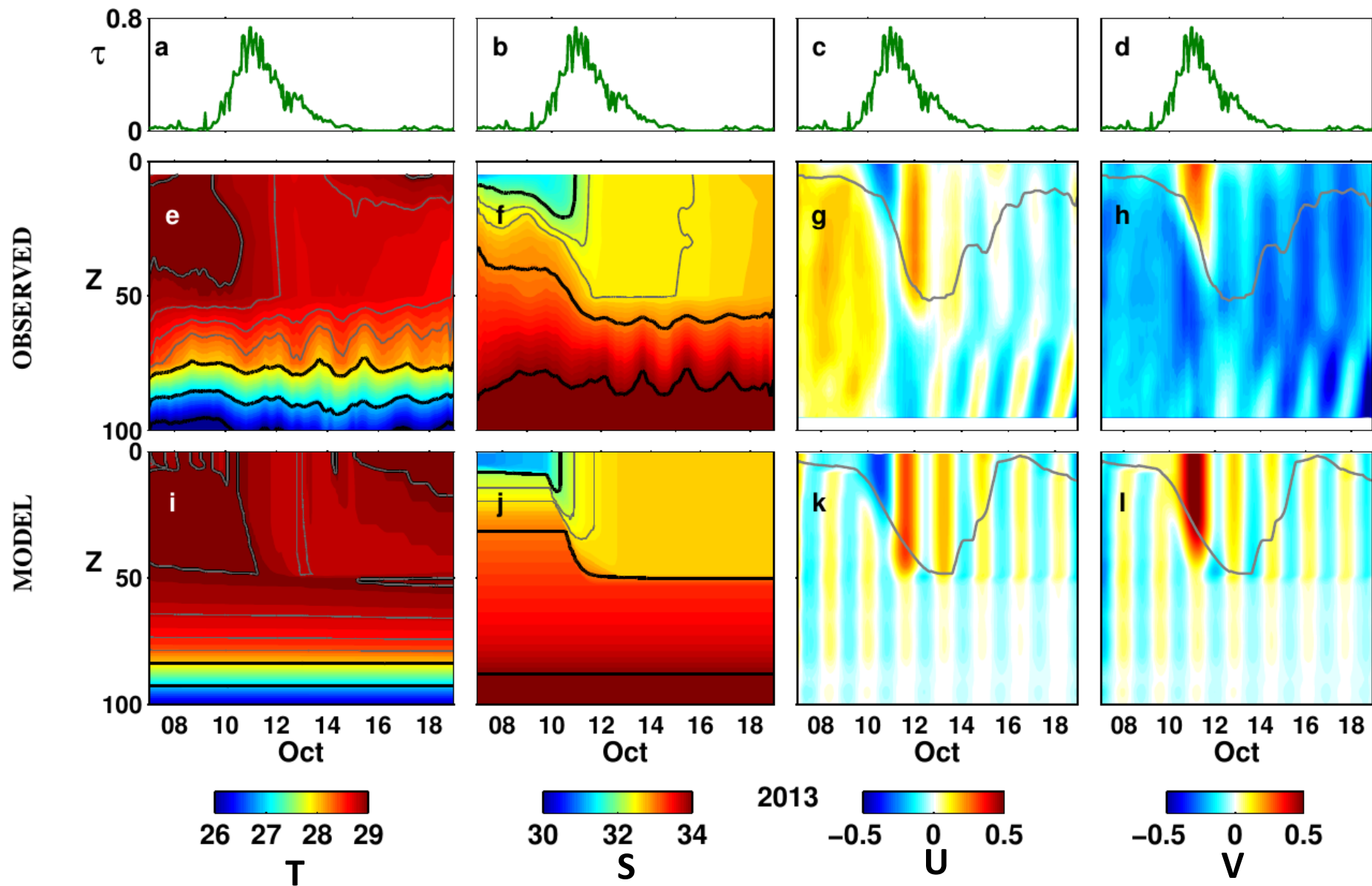
$$\text{Ekman depth} = 2.2 / 0.19 = 11.5 \text{ m}$$

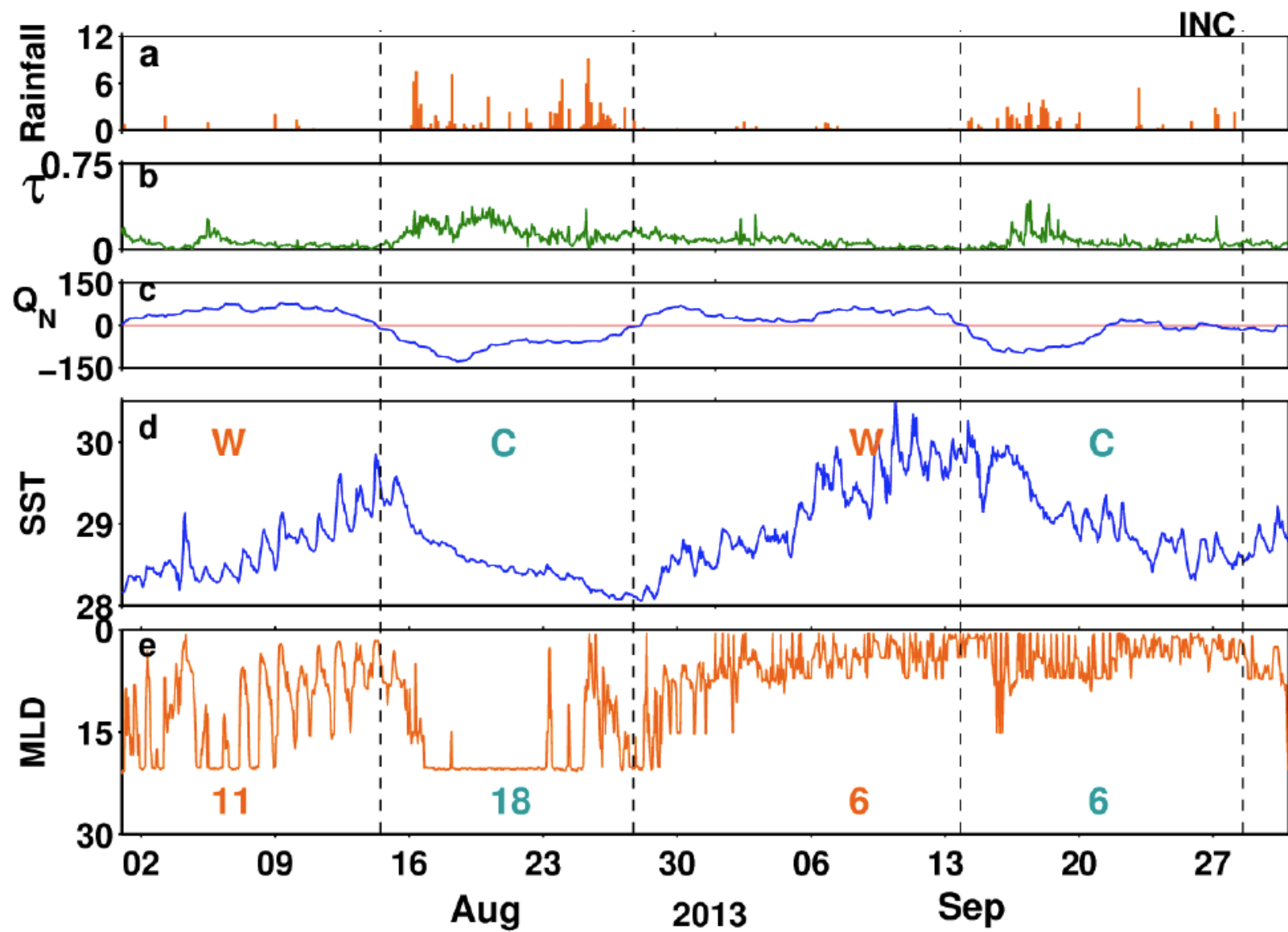


BD09

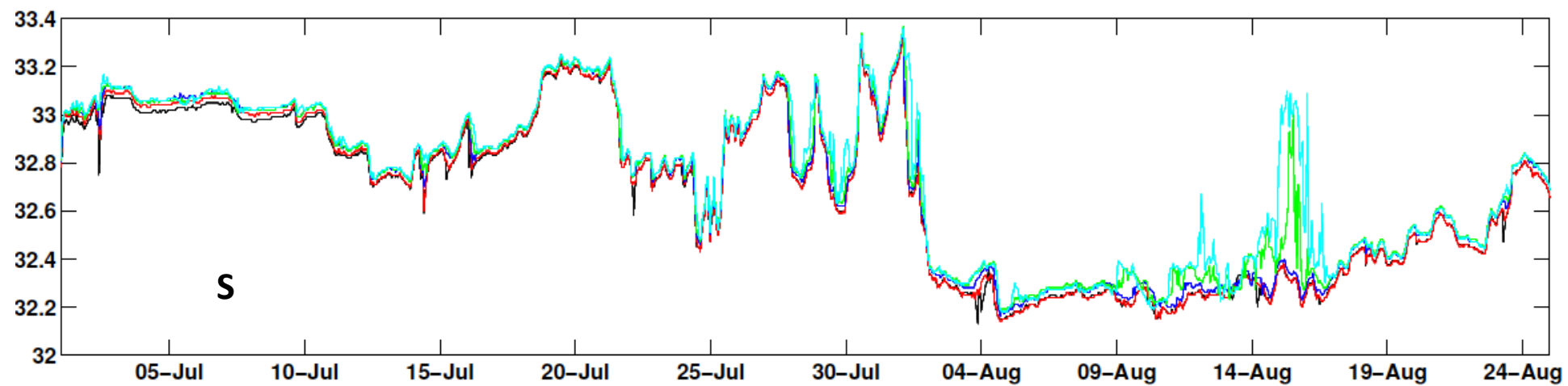
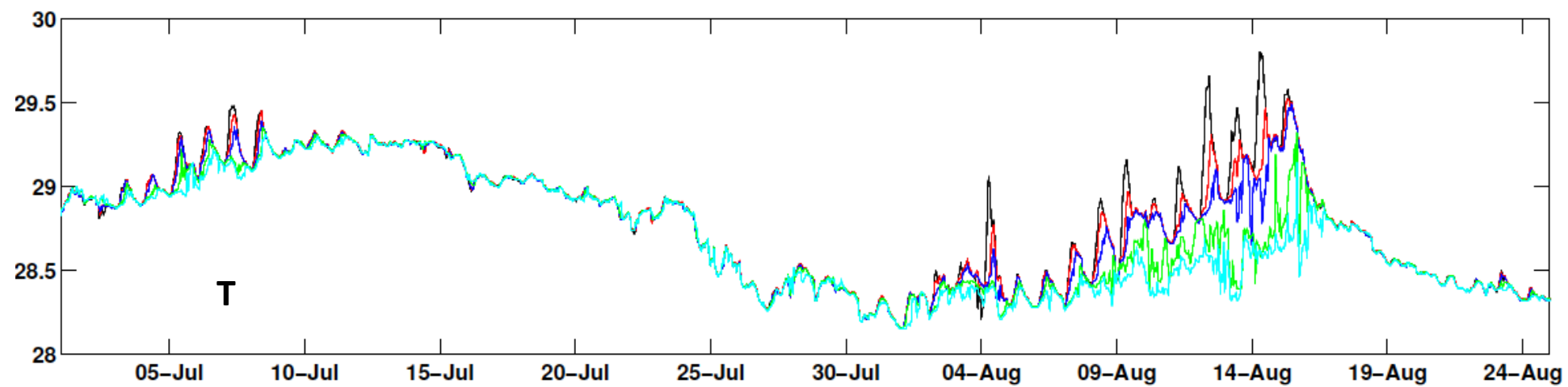


# PWP Model

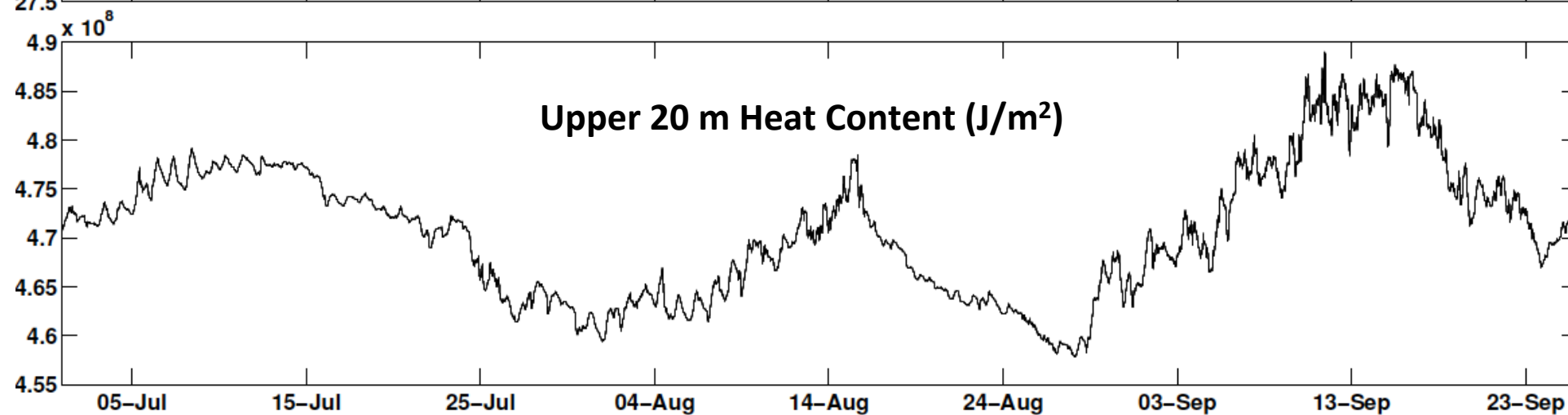
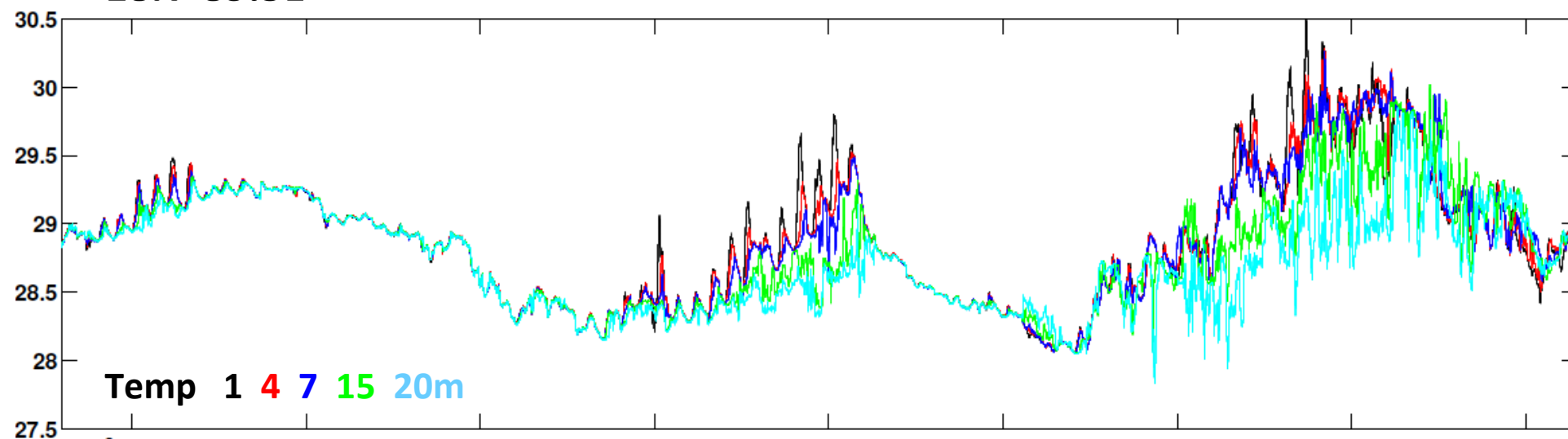


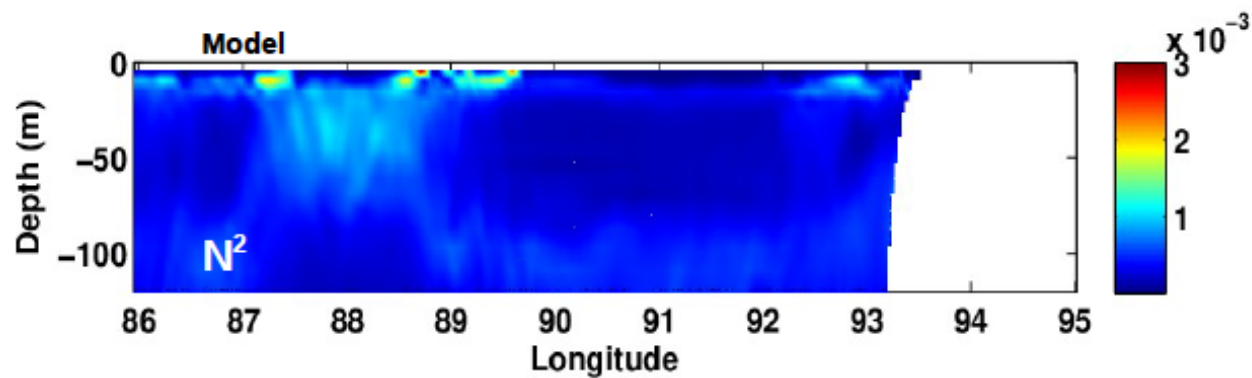






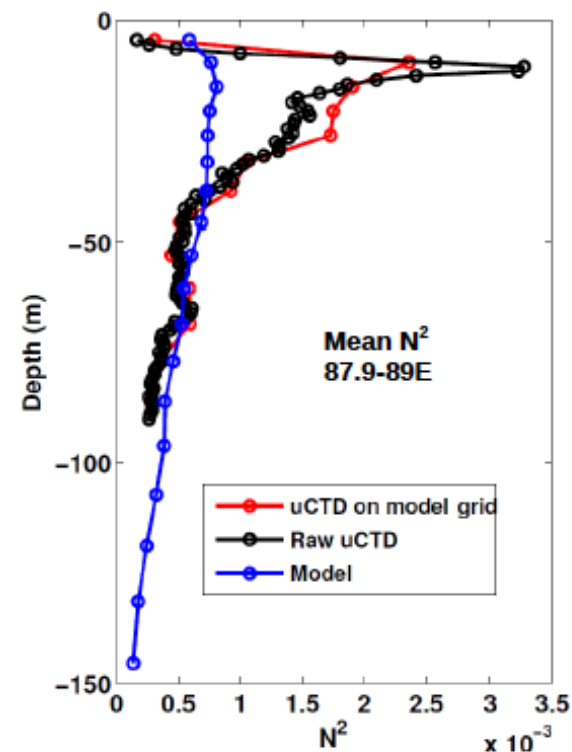
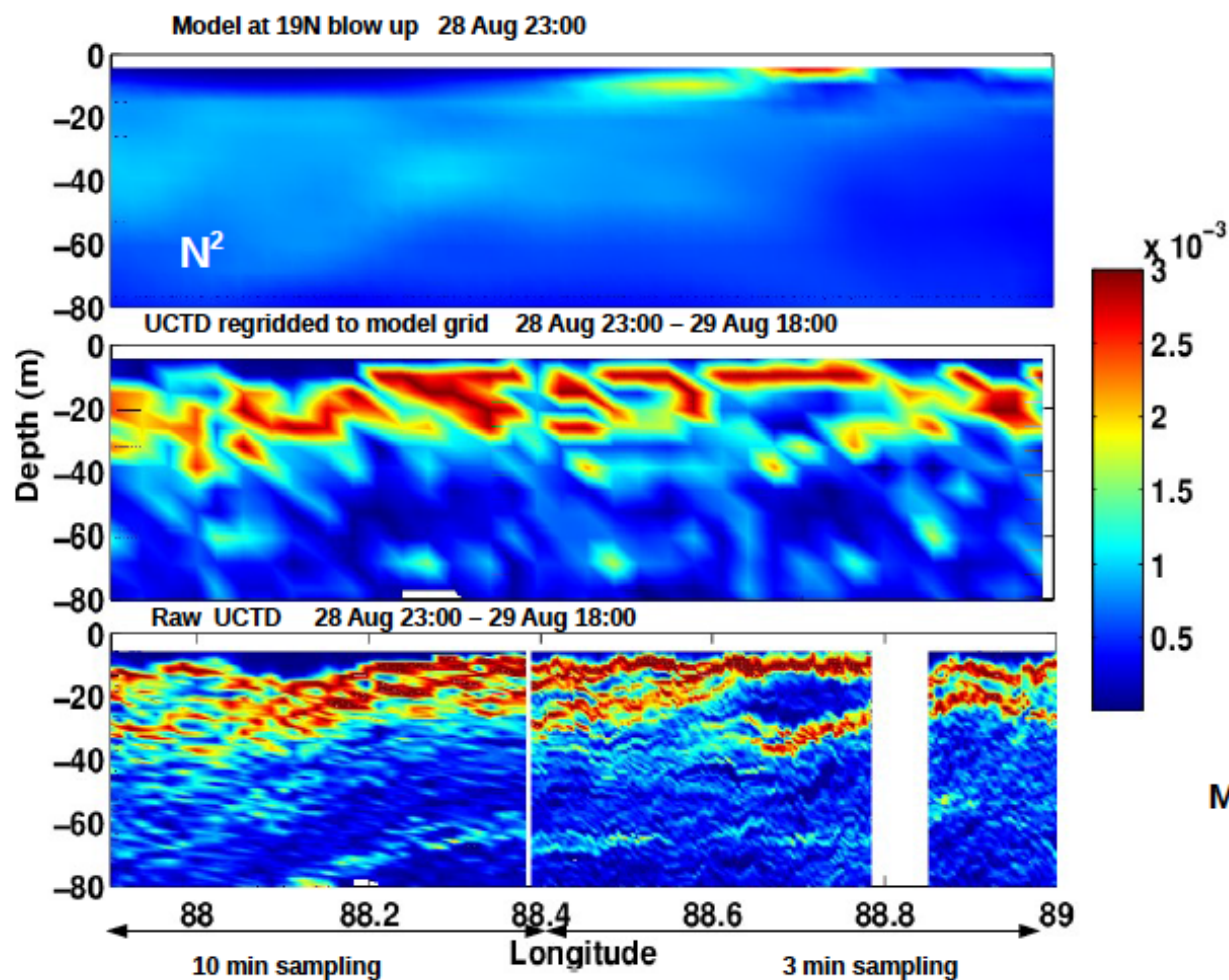
**18N 89.5E**





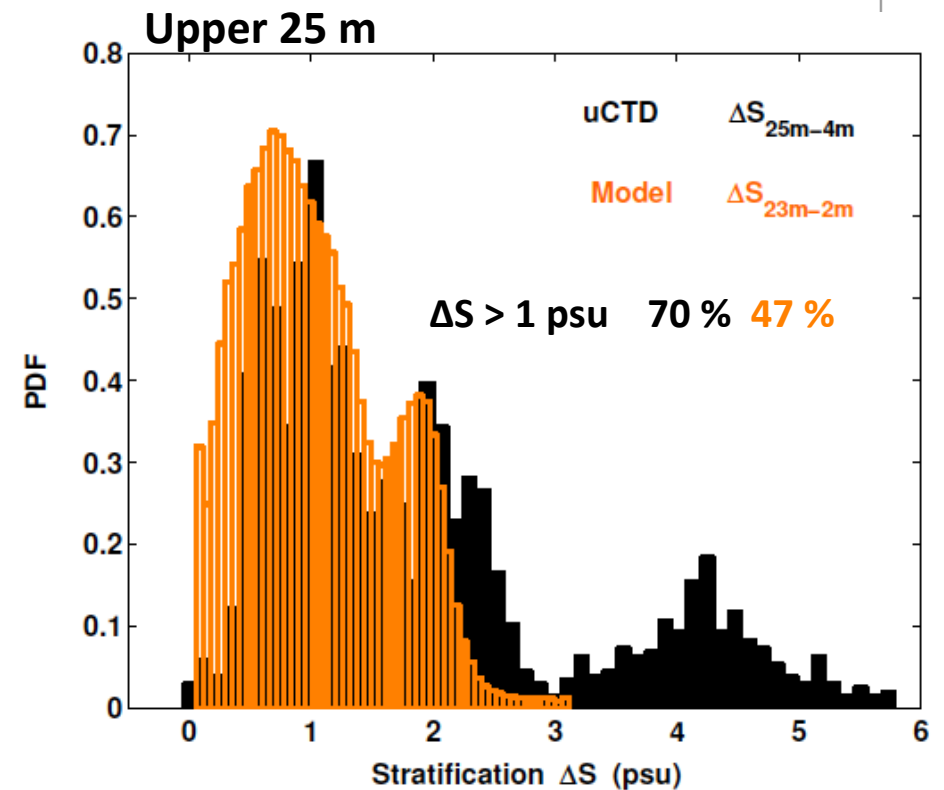
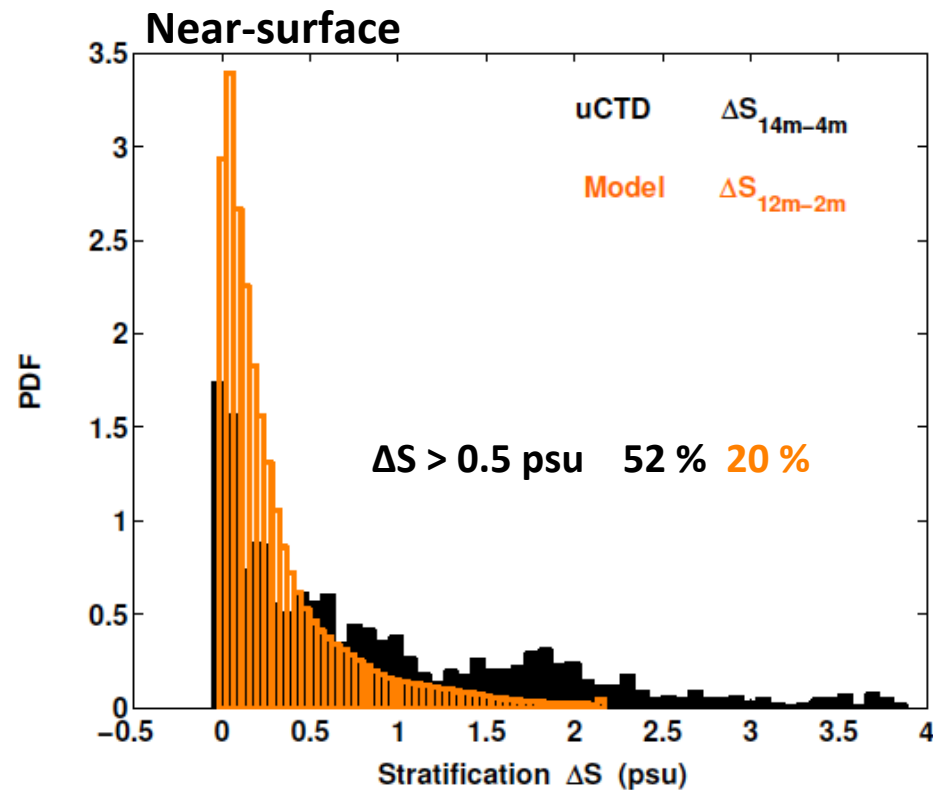
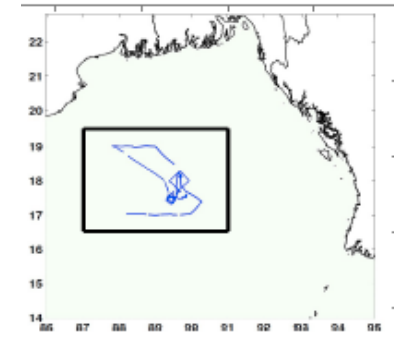
Model  $N^2$  at 19N 28 Aug 23:00

uCTD  $N^2$  at 19N  
28 Aug 23:00 – 29 Aug 18:00



Model stratification is weak

## PDFs of observed and model vertical salinity gradients



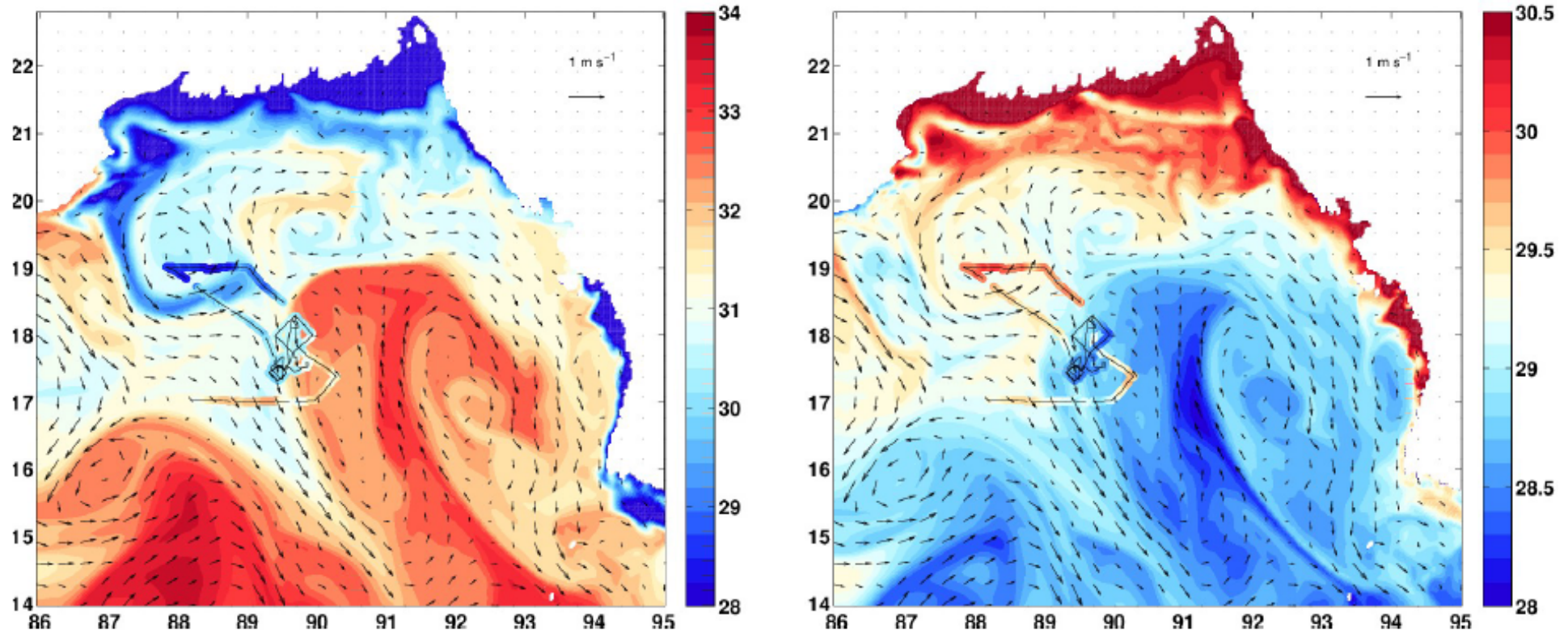
Model underestimates salinity stratification in fresh water



# 3 km ROMS 16 Aug-6 Sep 2014

Model 2 m Salinity (left), Temperature (right) and currents

26 Aug 06:00 – 27 Aug 06:00

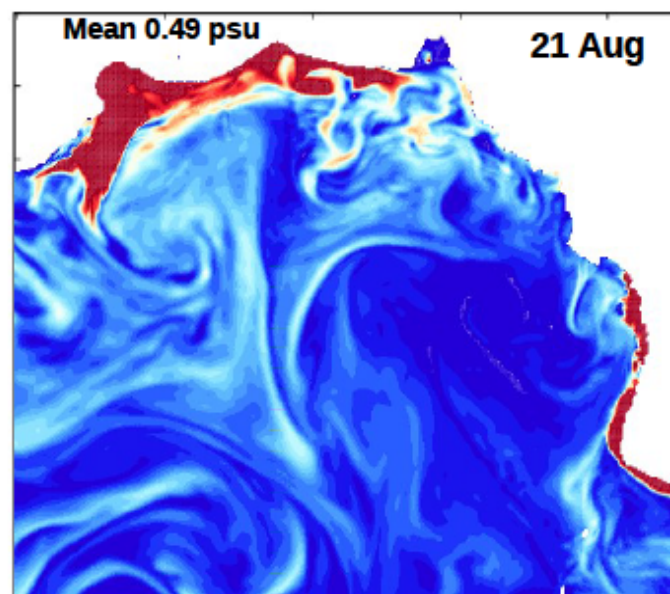
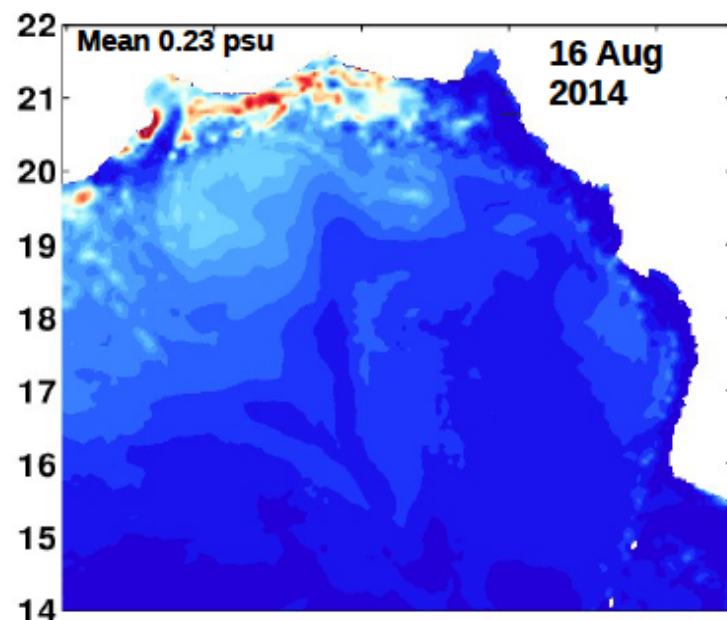


**Bay of Bengal Regional Ocean Model**

**1/32° 50 sigma levels Two-equation turbulence model**

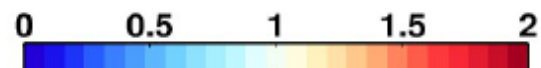
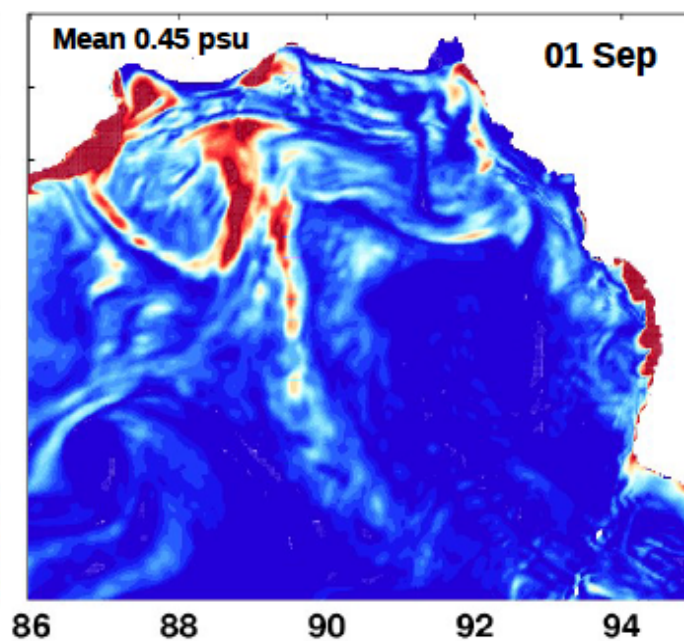
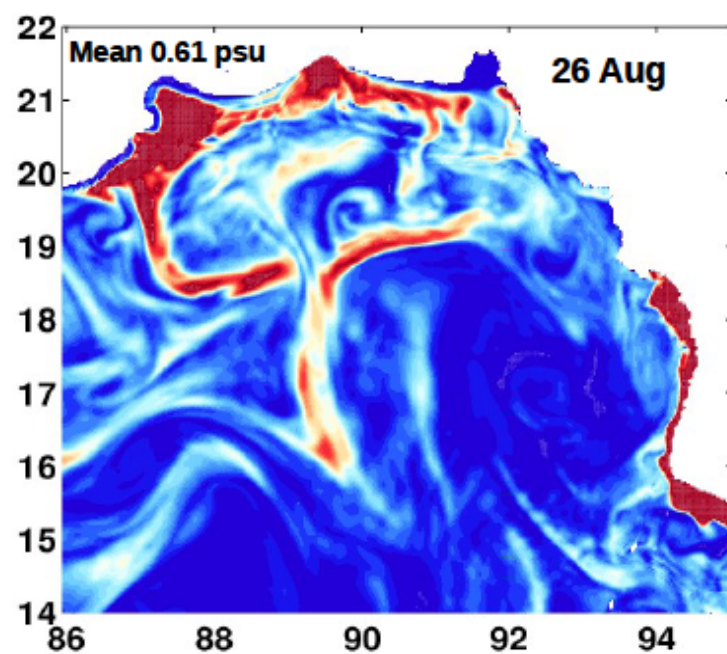
**Initialised from data-assimilating HYCOM on 16 August 2014**

**Forced with 3-hourly MERRA surface fluxes, hourly model fields**

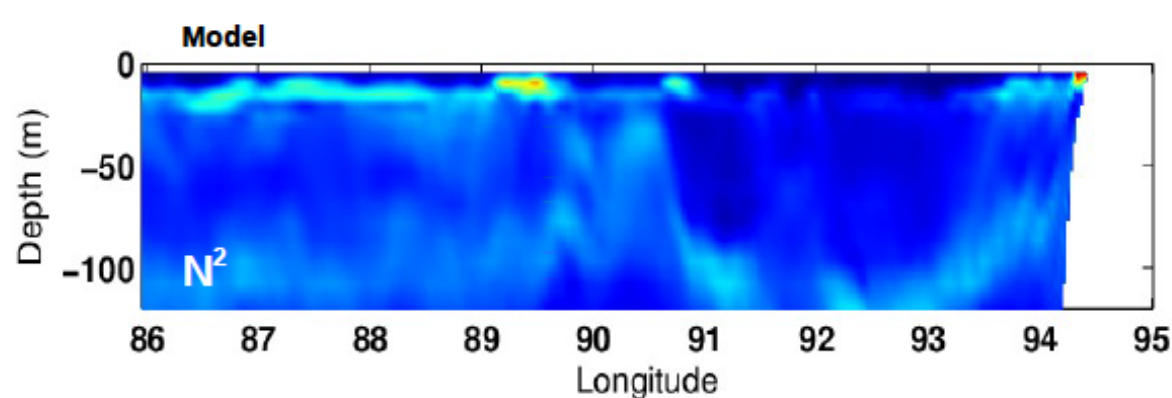


Model upper 10 m  
stratification

$\Delta S_{12m-2m}$  psu

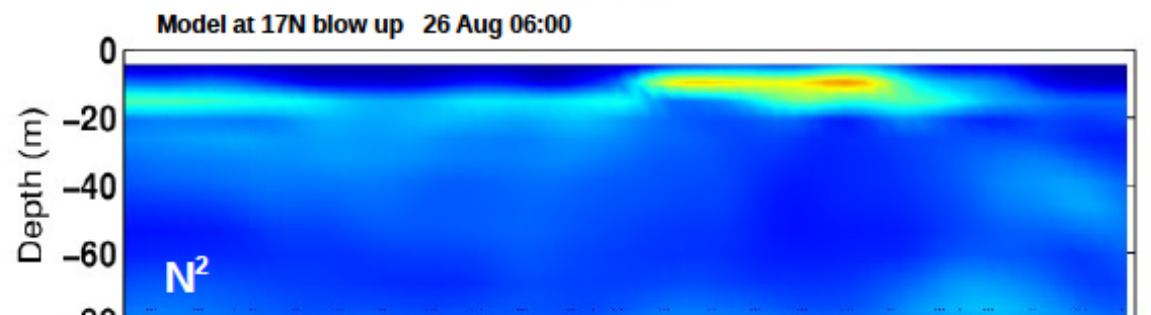




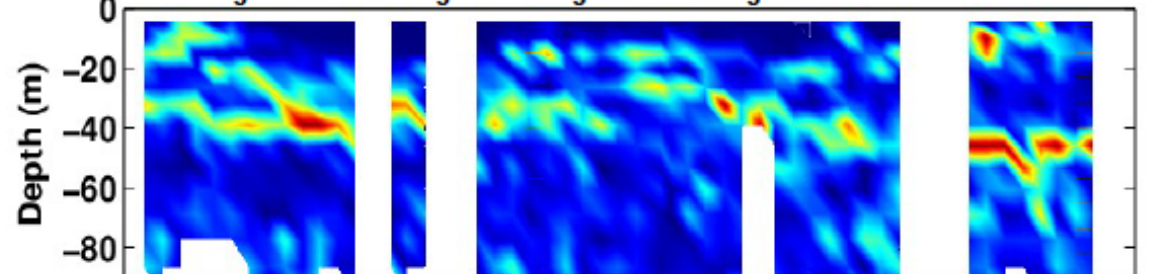


**Model  $N^2$  at 17N 26 Aug 06:00**

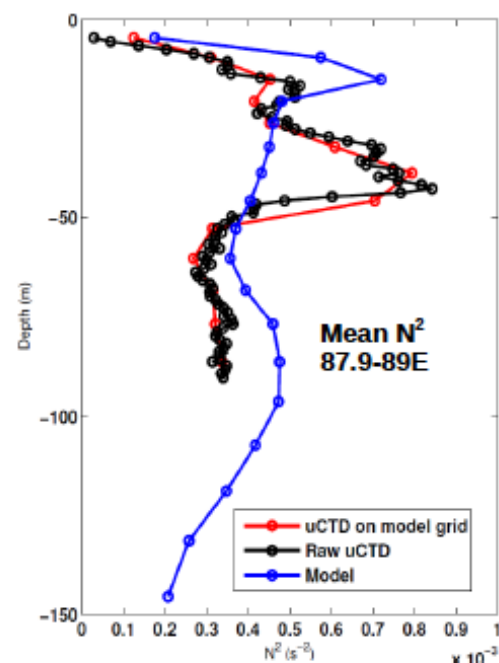
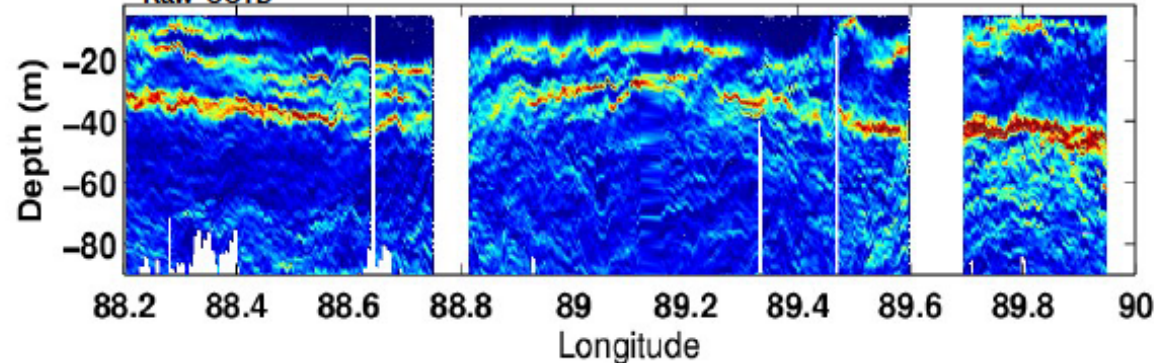
**uCTD  $N^2$  at 17N  
26 Aug 06:00 – 27 Aug 06:00**



**UCTD regridded to model grid 26 Aug 06:00 – 27 Aug 06:00**

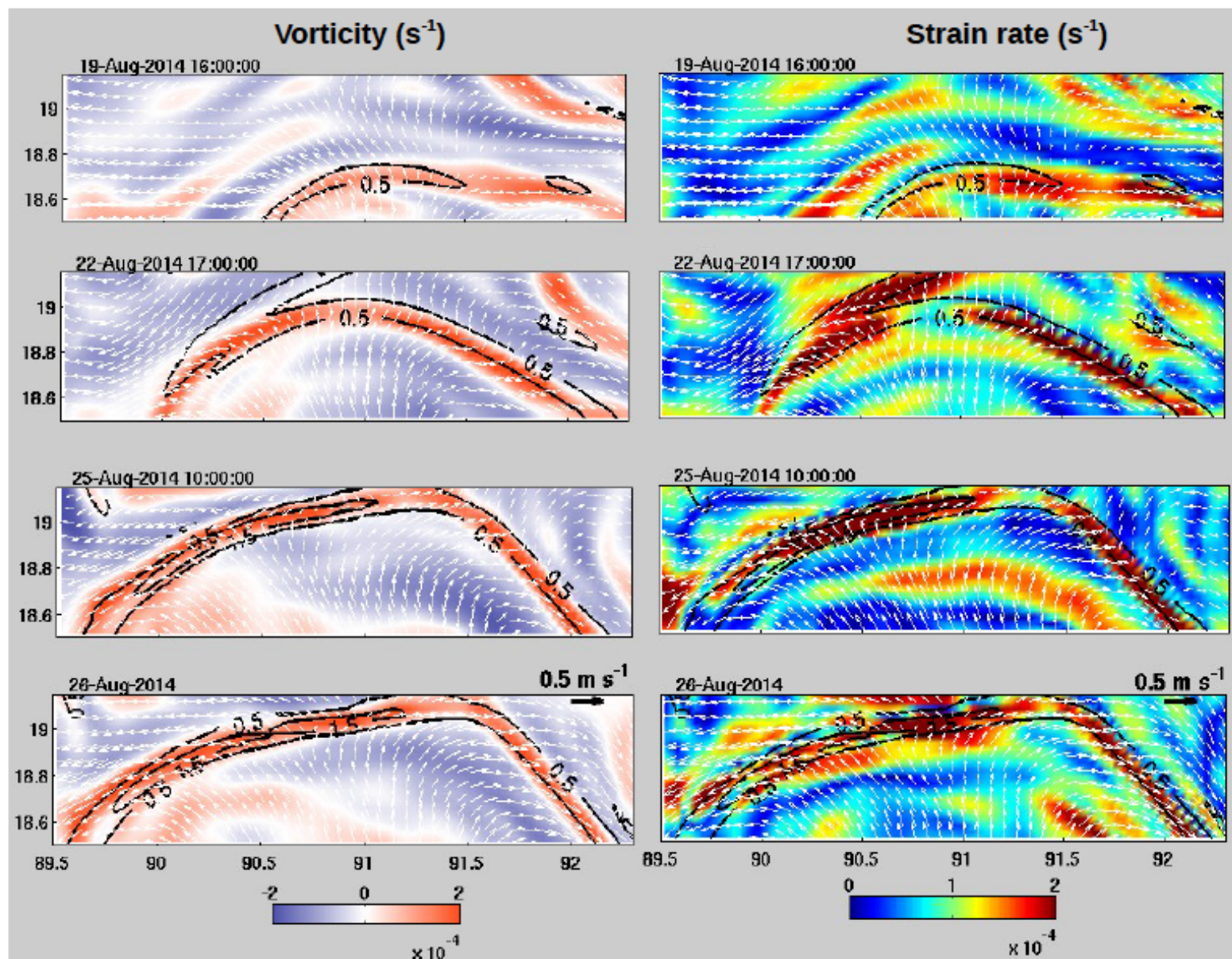


**Raw UCTD**



**uCTD  $N^2$  max at 30-50 m.**

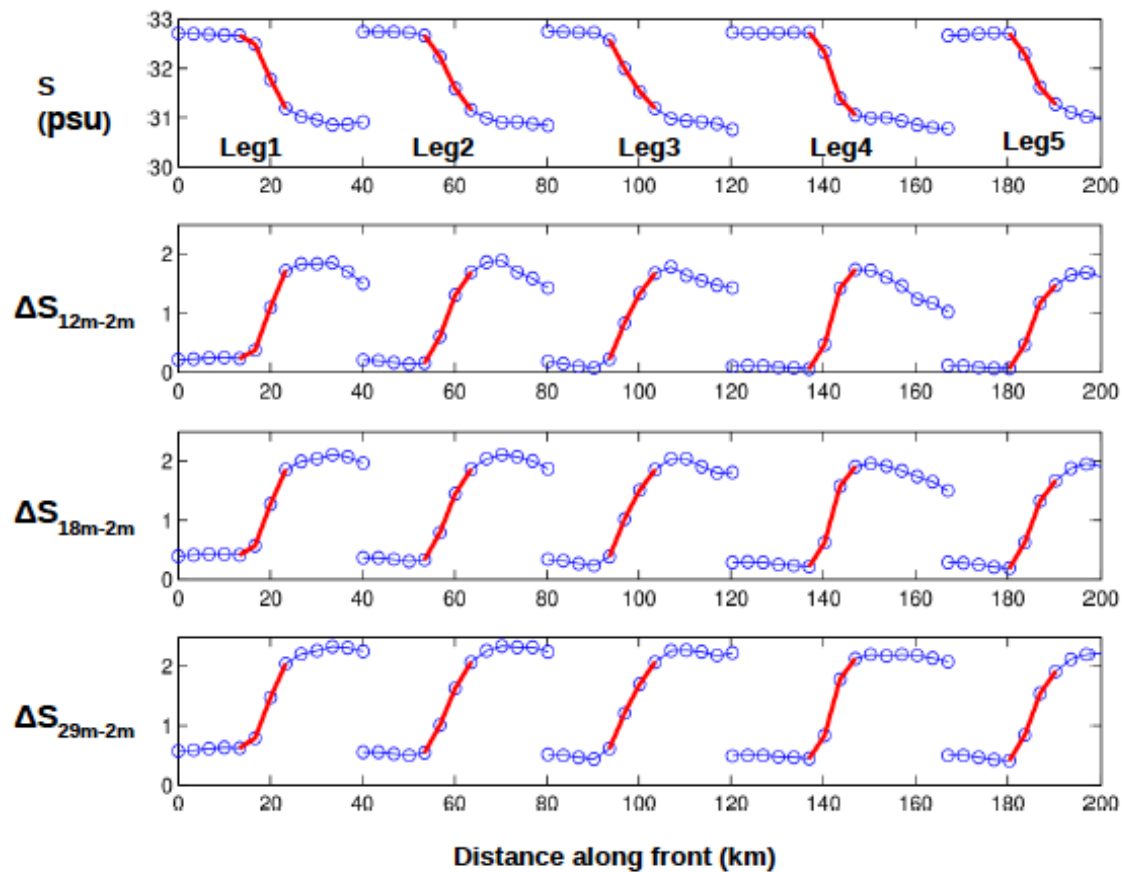
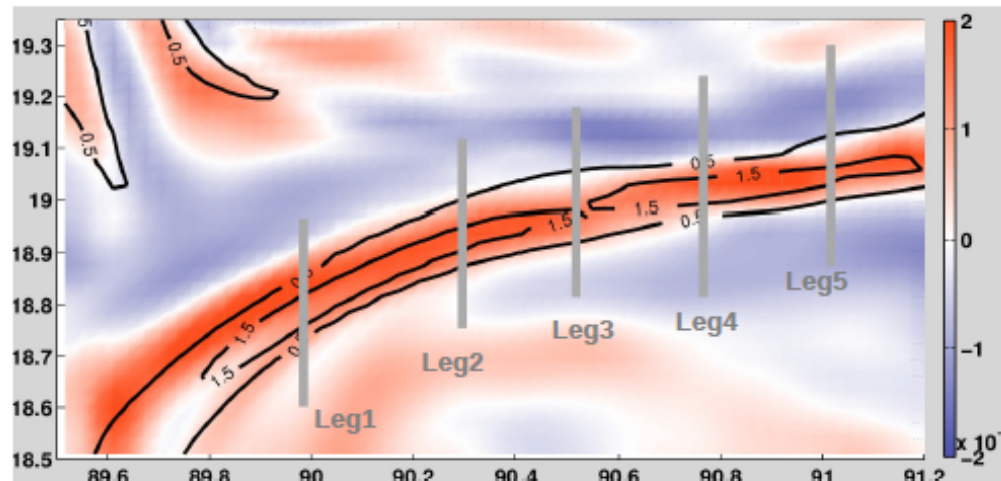
**Model  $N^2$  max at 5-20 m.**



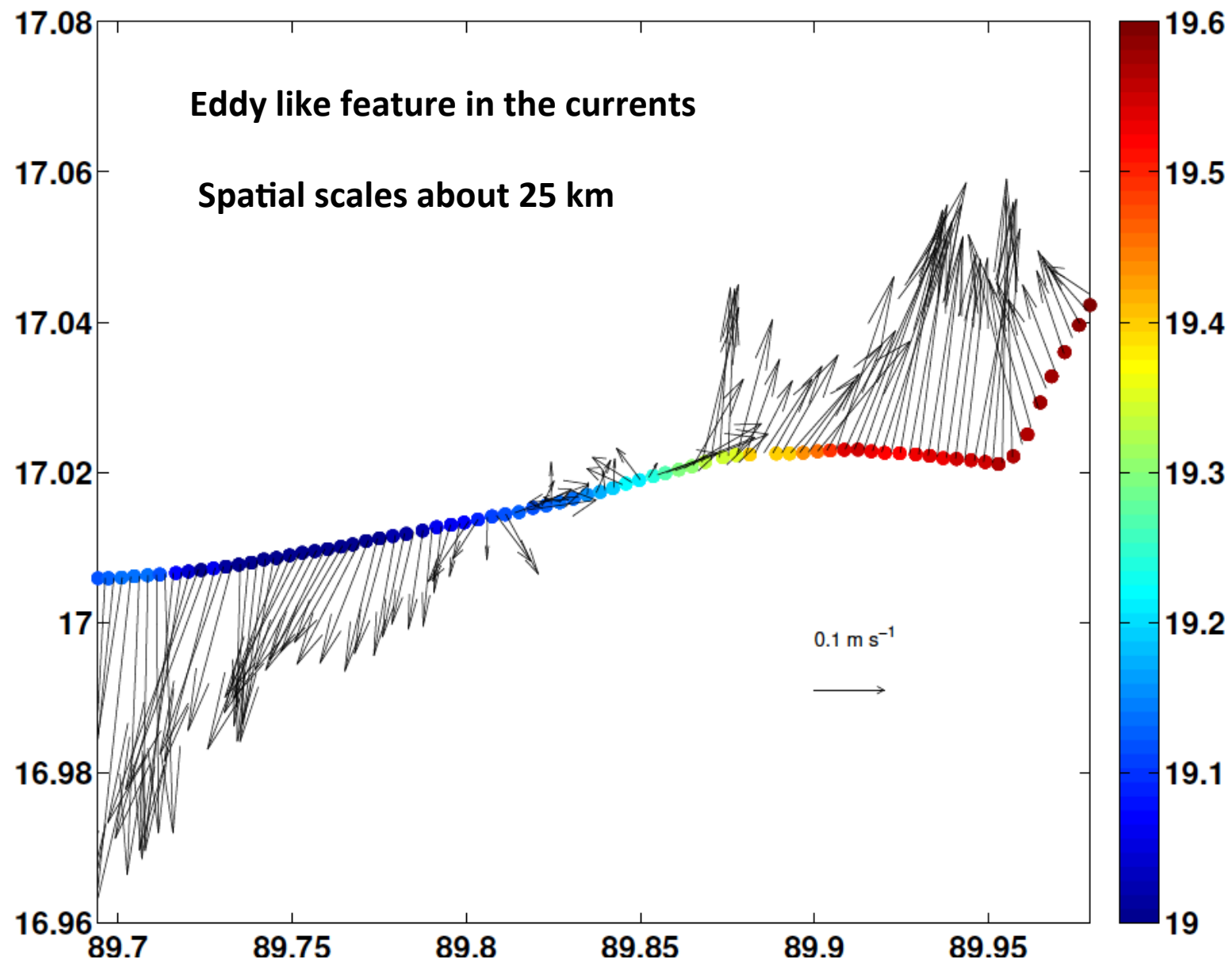
$|\text{grad}S_{x,y}|$  (psu/km; black contour)

Vorticity and strain rates are high at the front





The near-surface salinity stratification is higher on the fresher side of the front rather than directly under the front



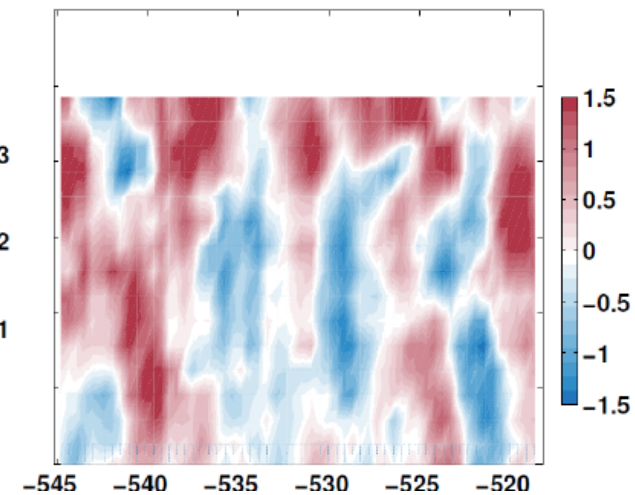
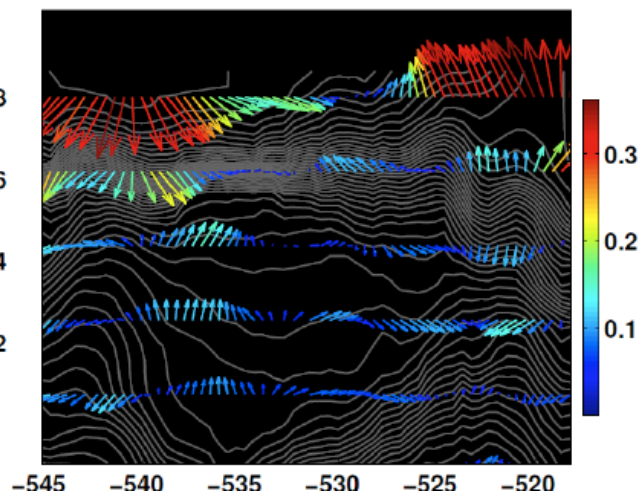
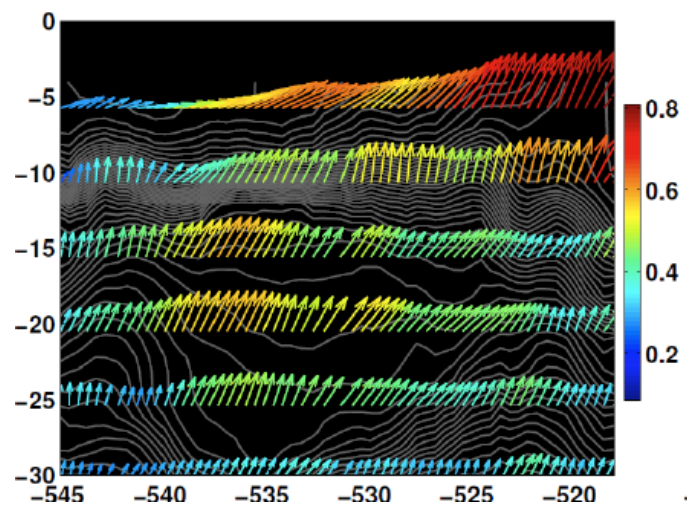
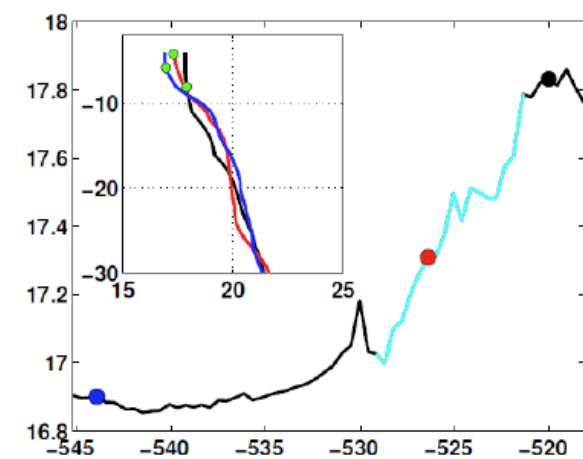
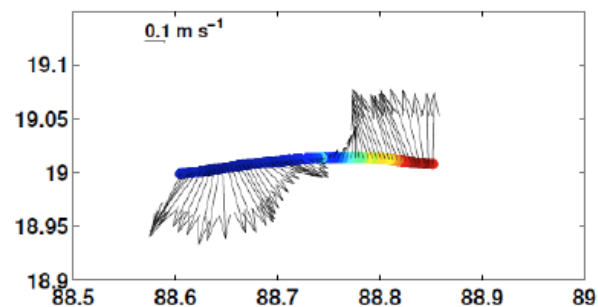
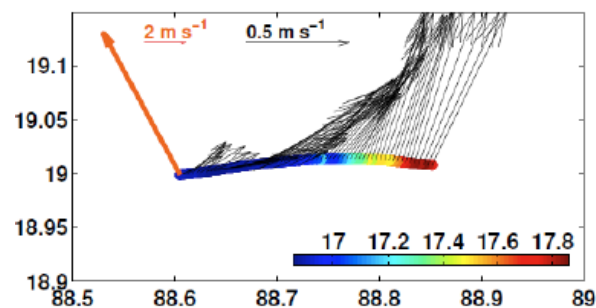


[illegible]

***D'Anville 1762***

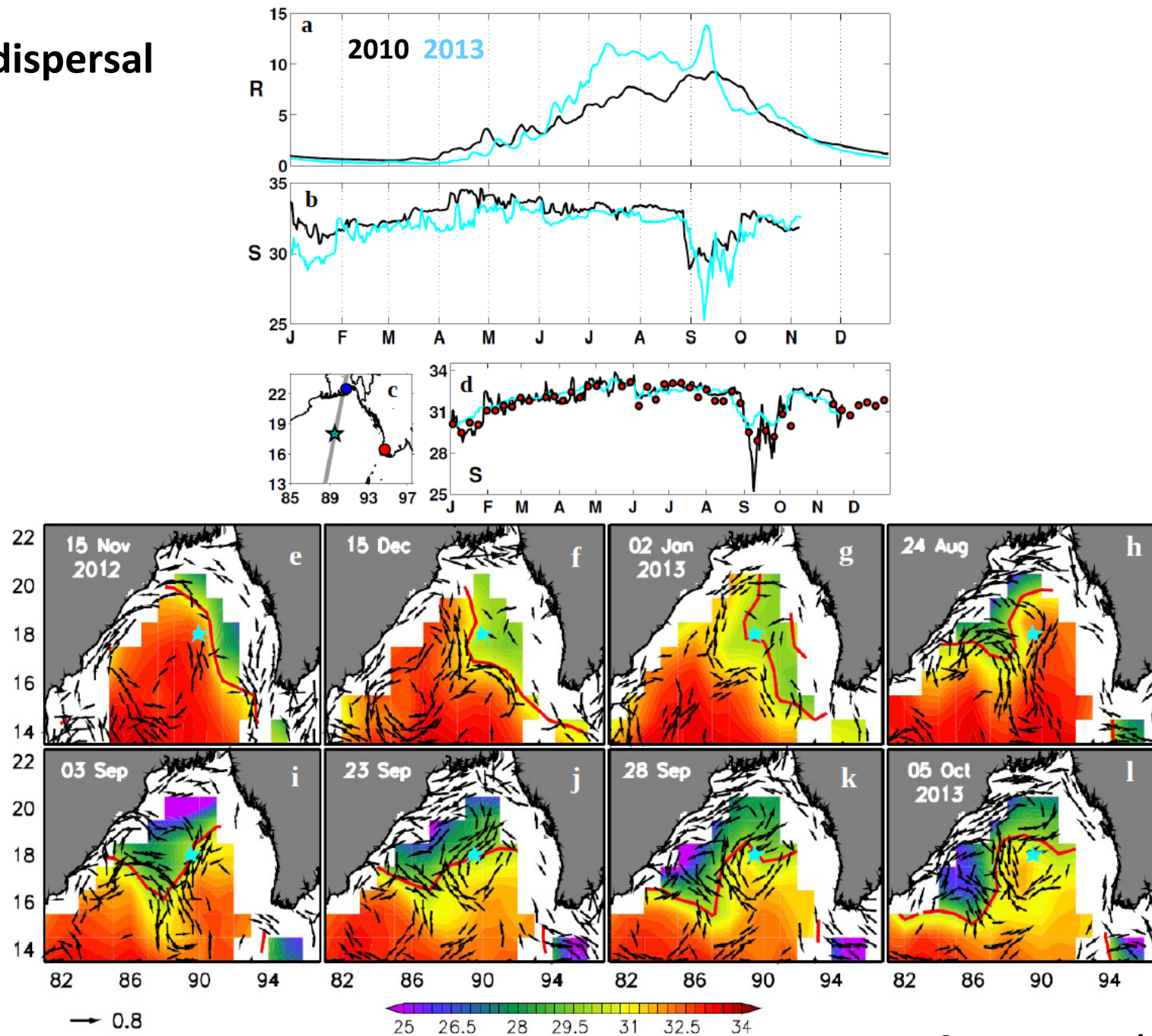


## Submesoscale cyclonic eddy?

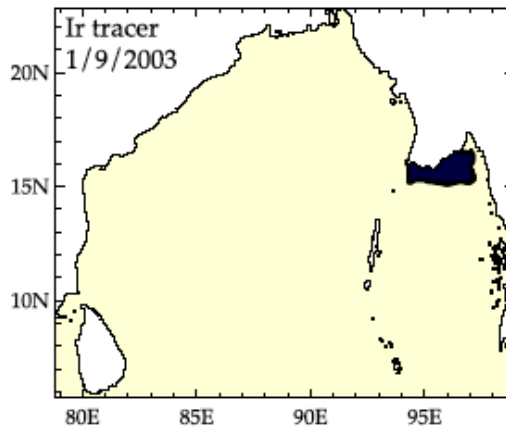
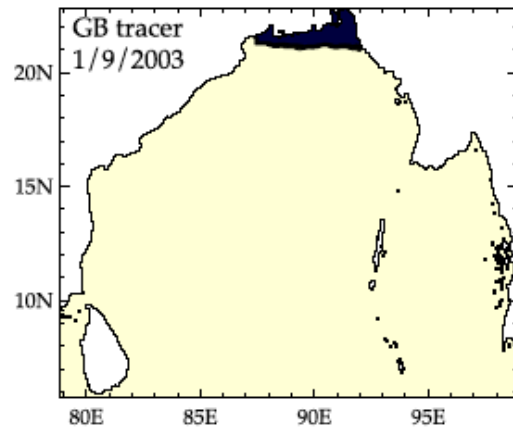




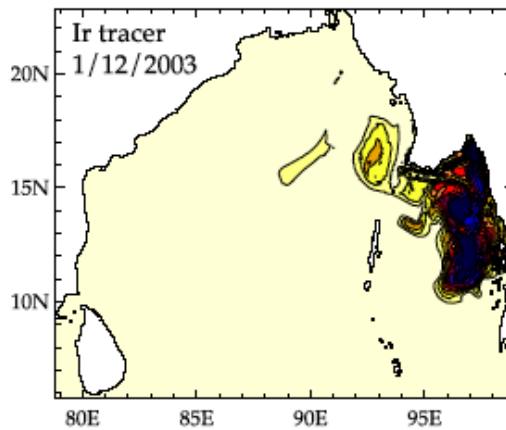
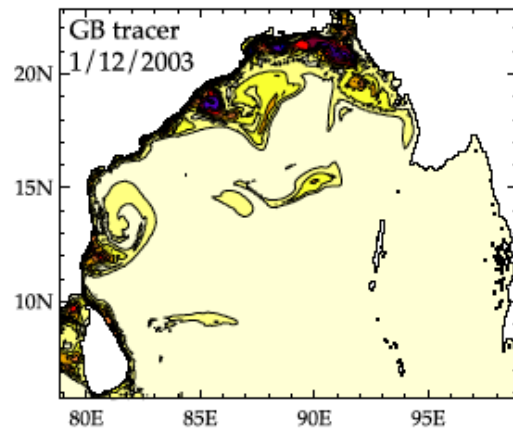
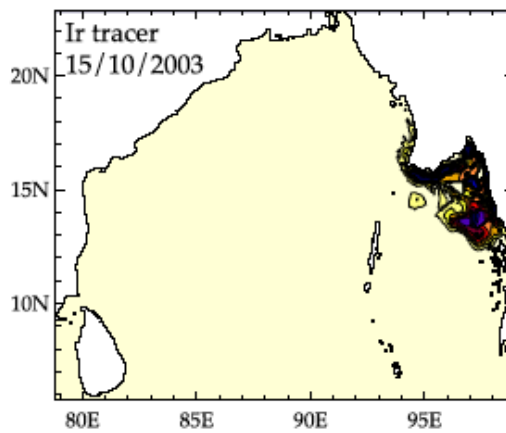
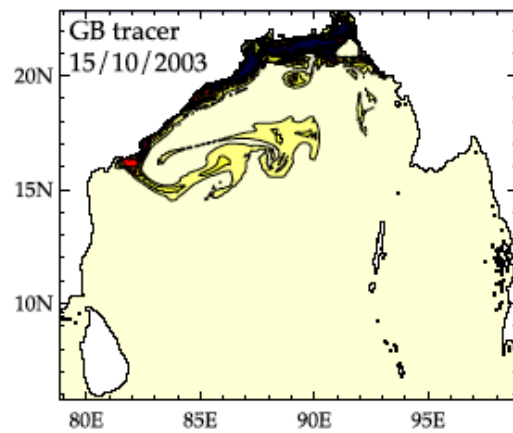
# FW dispersal



*Sengupta et al., 2016*



**NEMO 8 km 1 m res.**



**Benshila *et.al* 2014**

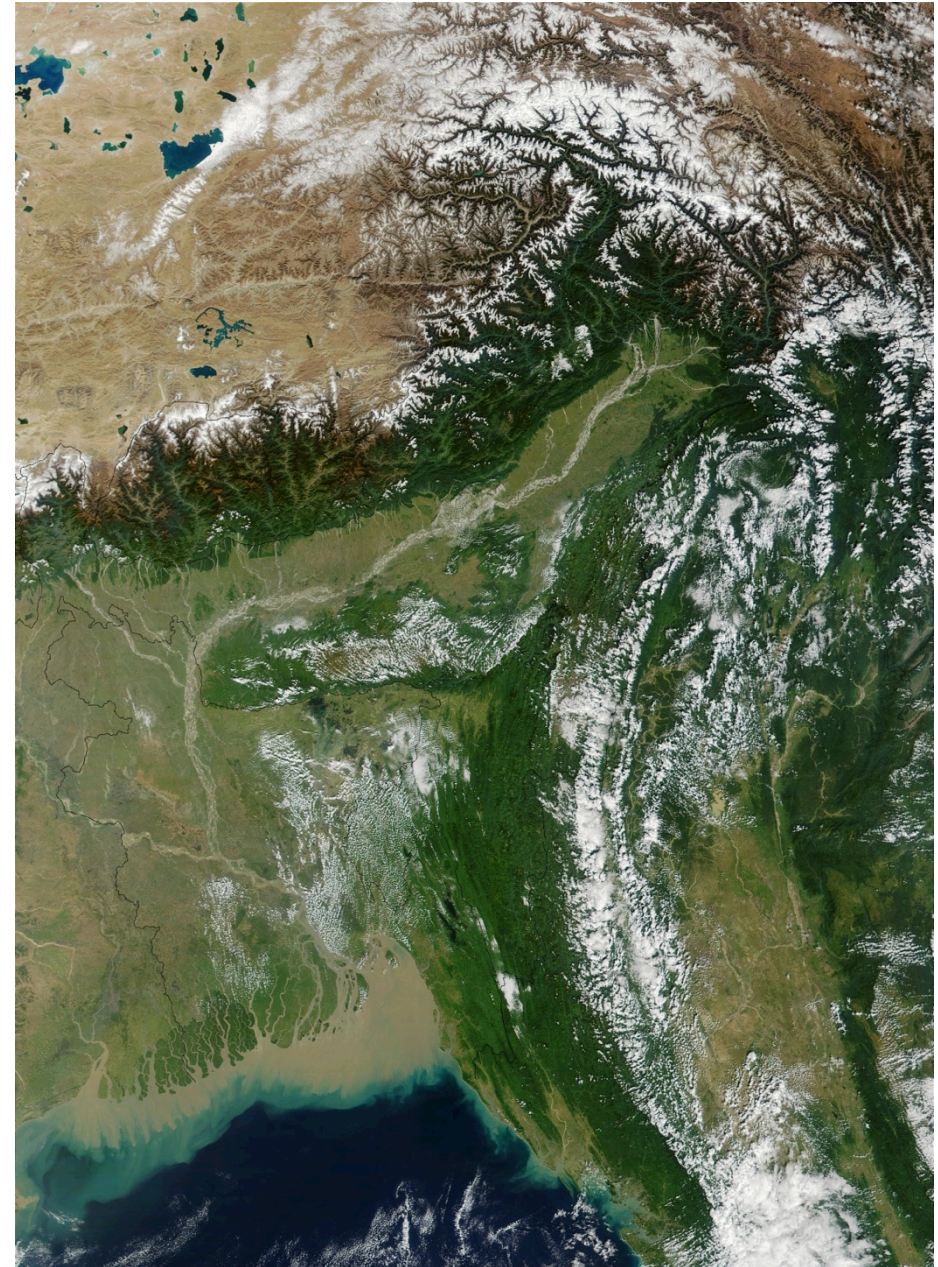


**Total freshwater input to the Bay of Bengal (P+R-E) is 4200 km<sup>3</sup> or 1.6 m**

**Several major monsoonal rivers flow into the north Bay of Bengal –**

**Ganga-Brahmaputra-Meghna,  
Irrawaddy, Salween, Godavari,  
Mahanadi, Krishna**

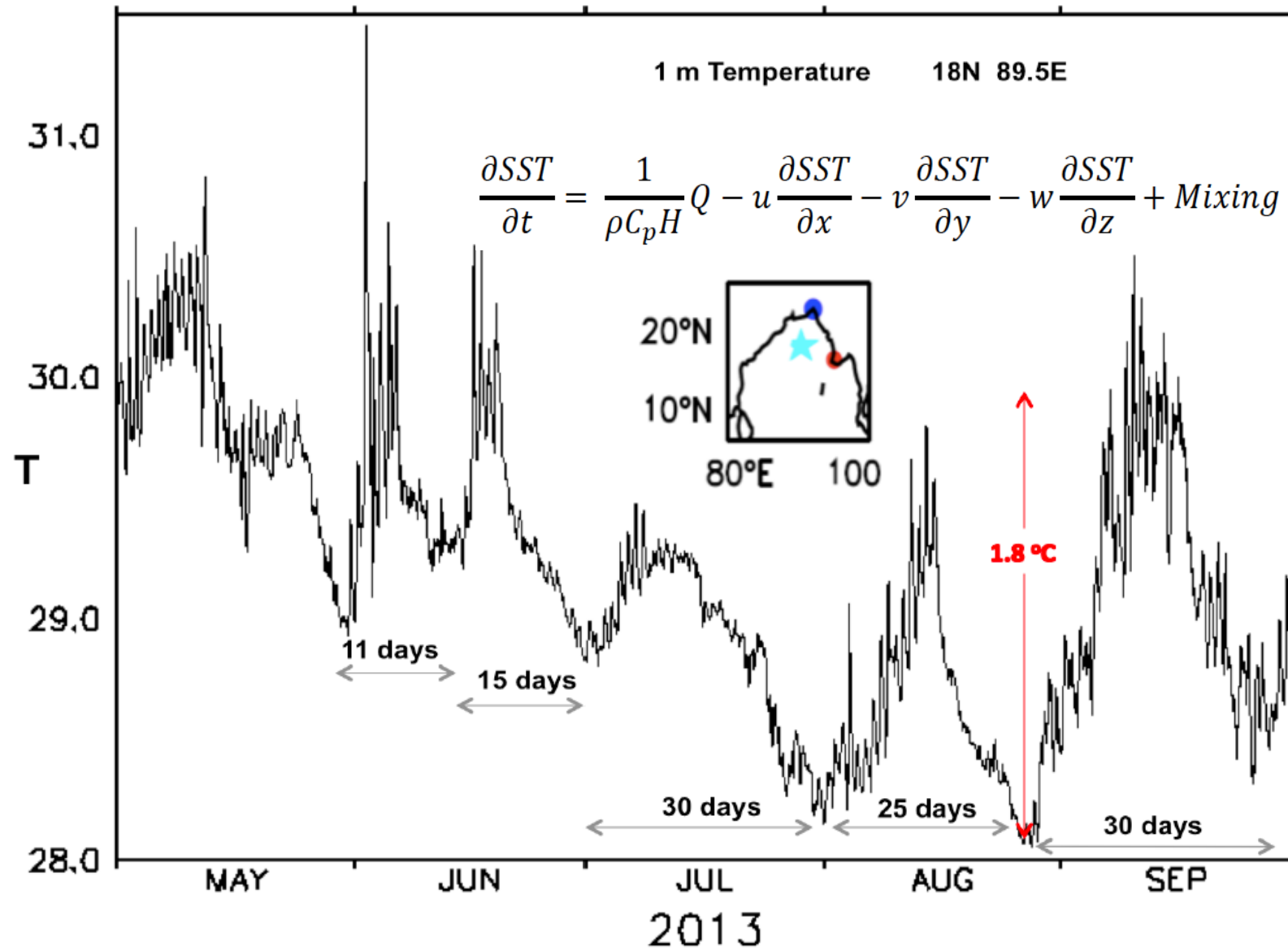
***Sengupta et.al 2006***



**Ganga-Brahmaputra-Meghna river**

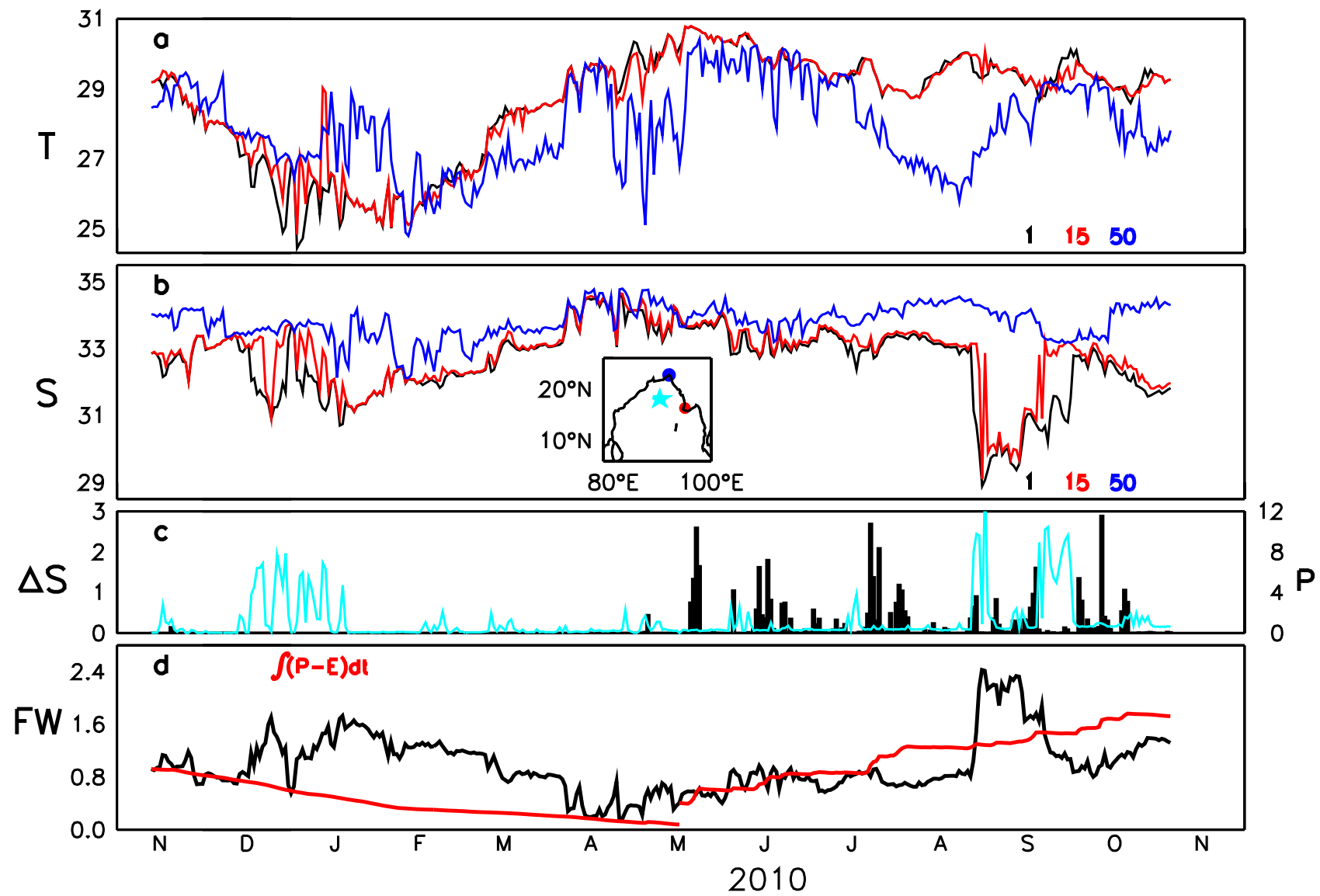
NASA

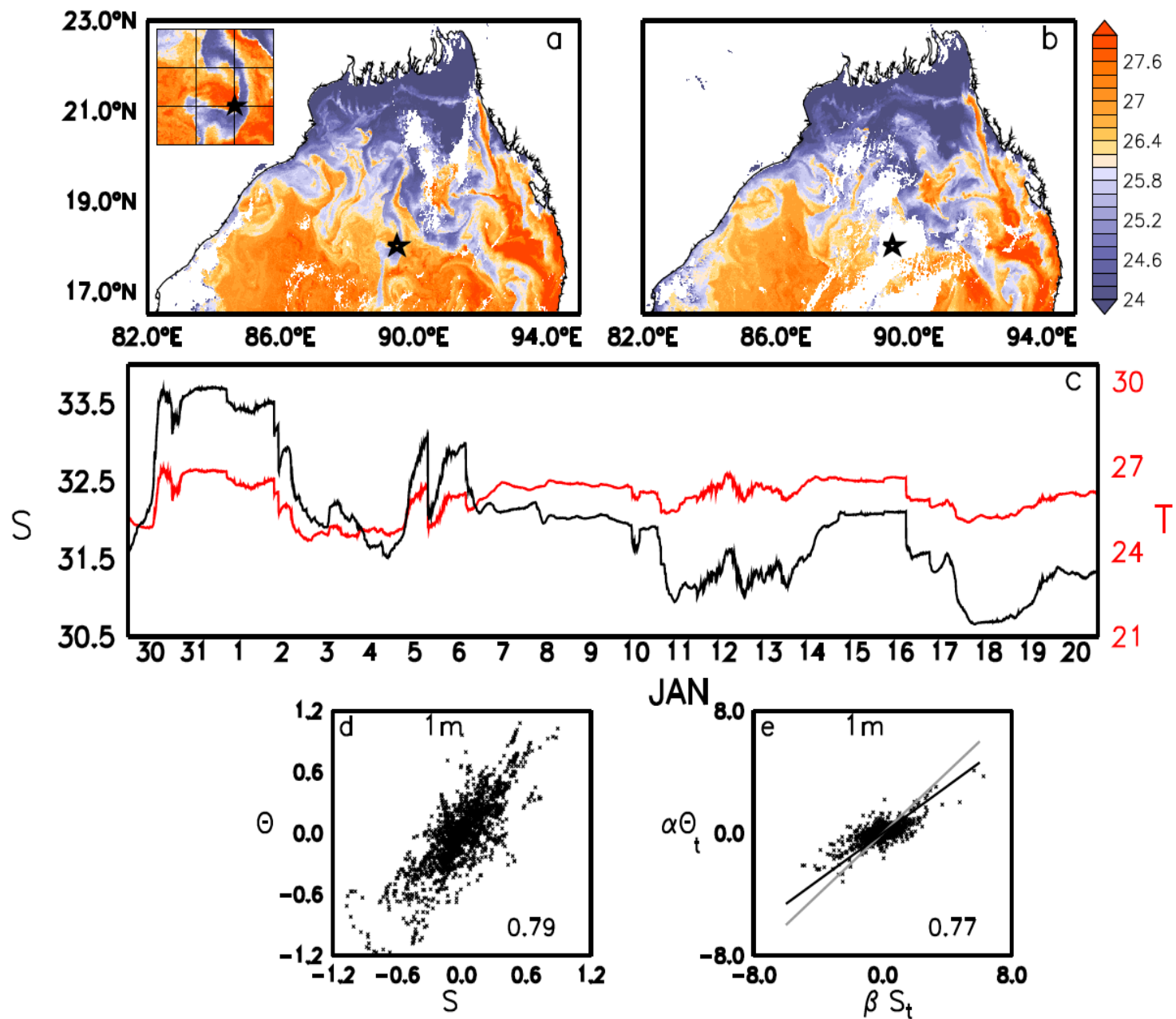
## Heat balance

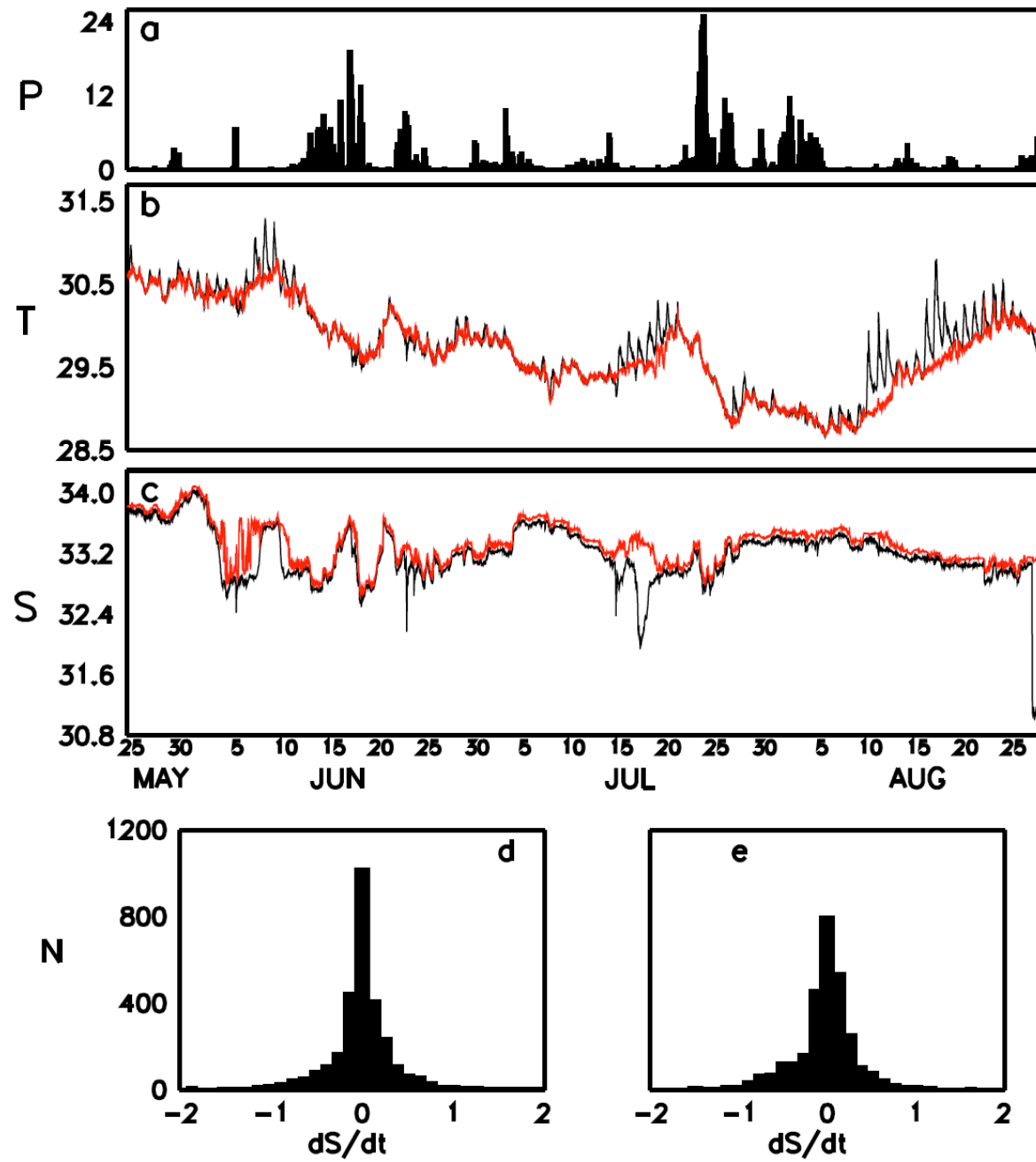


*Very shallow mixed layer*





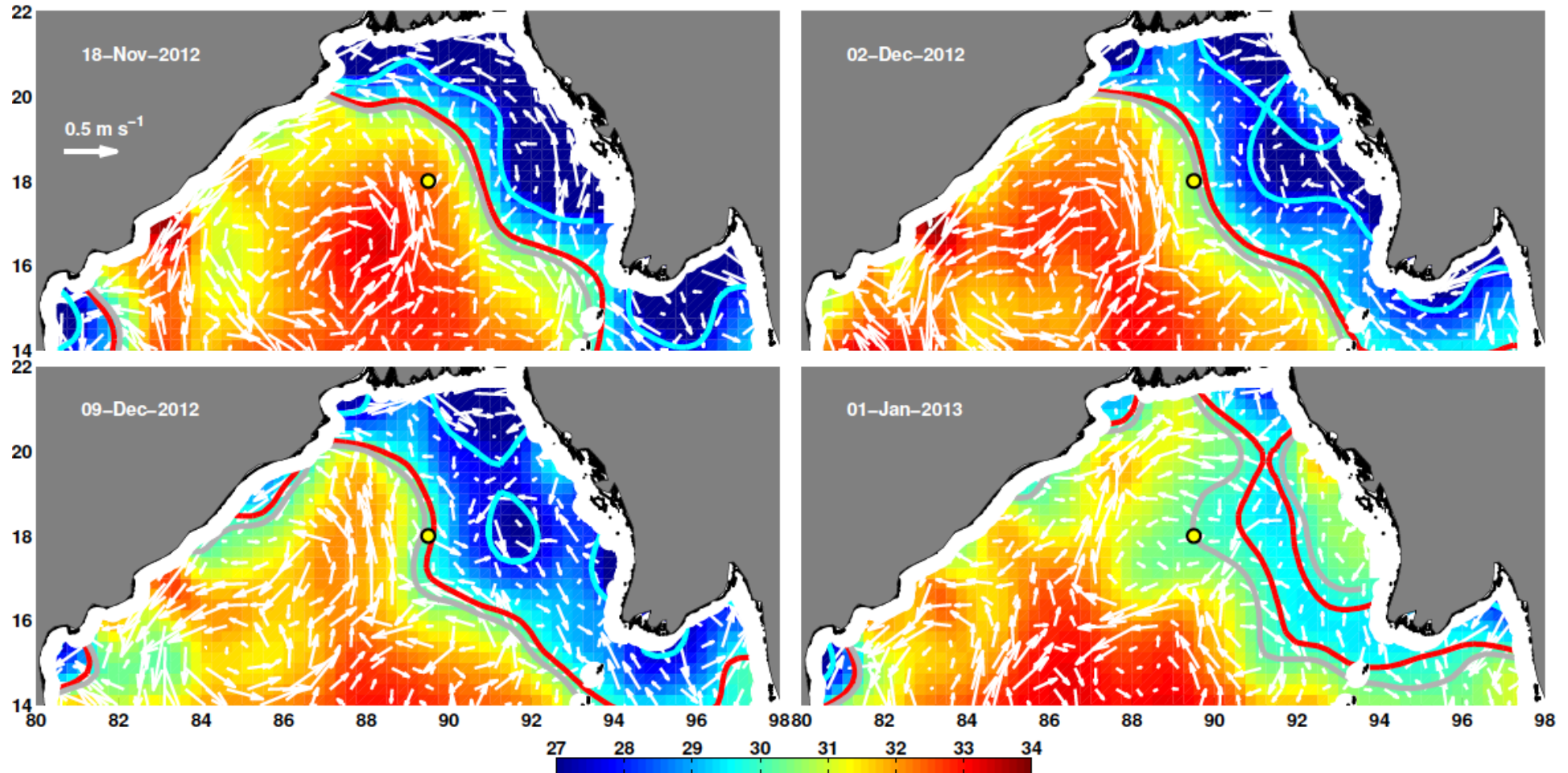




# Pathway Nov-Dec 2012

Aquarius sea surface salinity and AVISO geostrophic velocities

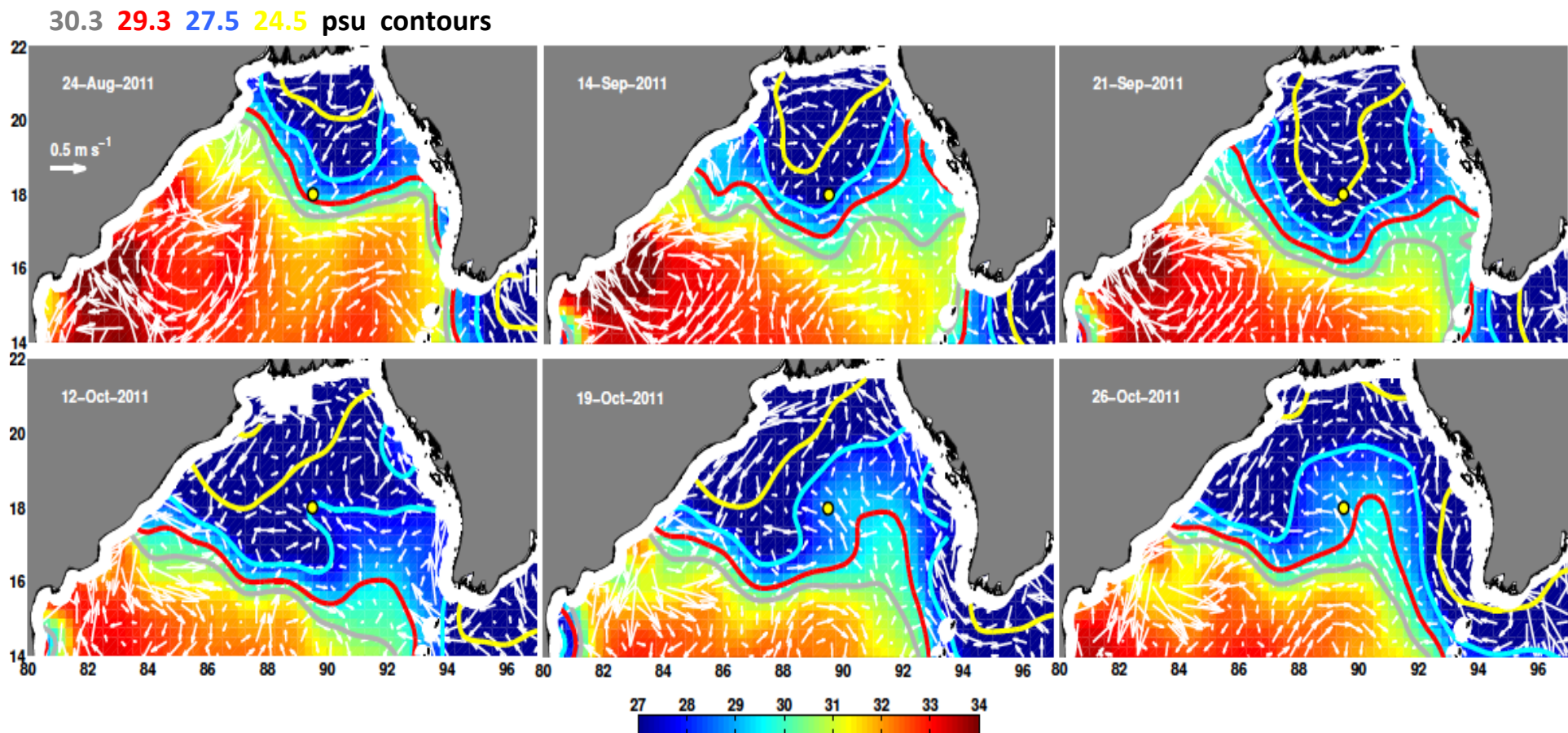
30.1 29.7 27.5 psu contours



Freshwater is advected by basin-scale flow and stirred by 200-400 km eddies.  
Winter pulse from Andaman sea, Irrawady water.



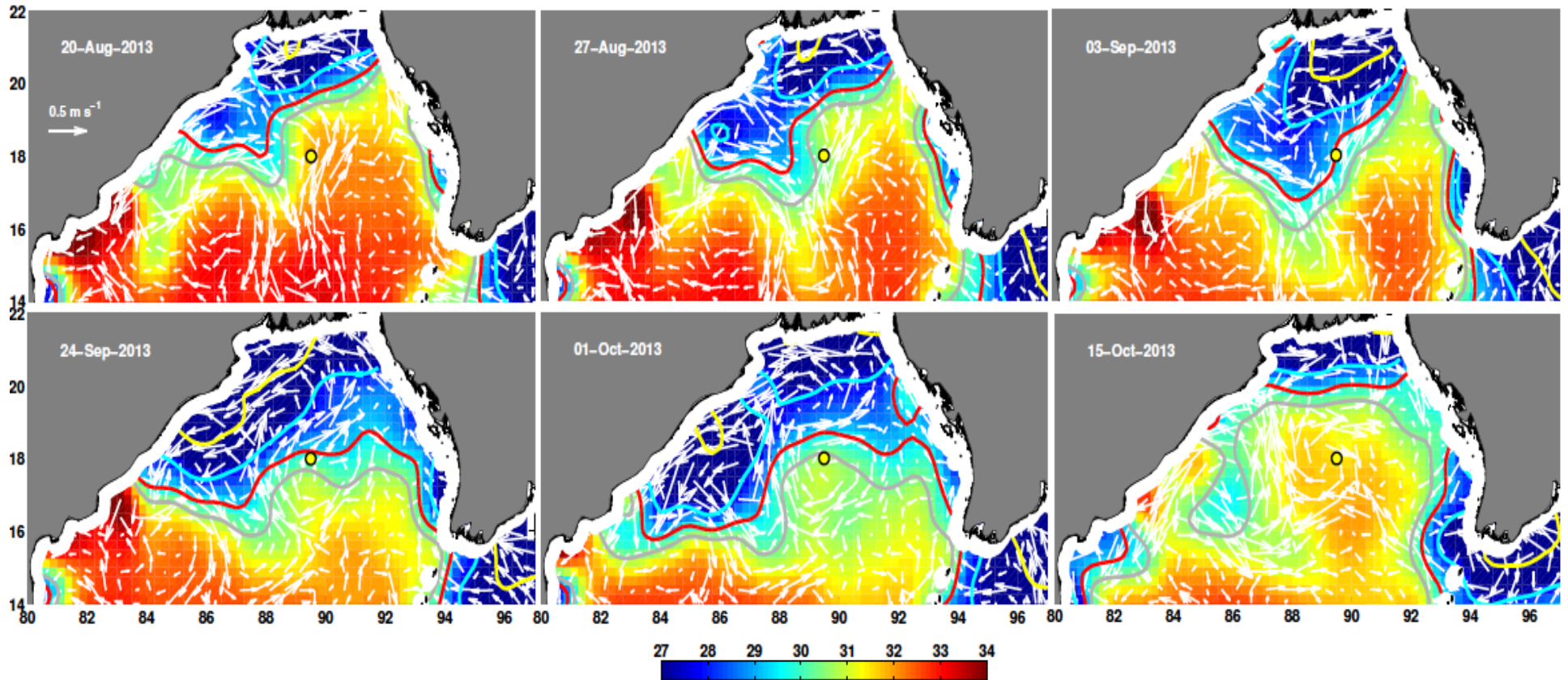
## Pathway Aug-Sep 2011



**Ganga-Brahmaputra-Meghna water flows down the center of the Bay.**

## Pathway Aug-Sep 2013

30.3 29.3 27.5 24.5 psu contours

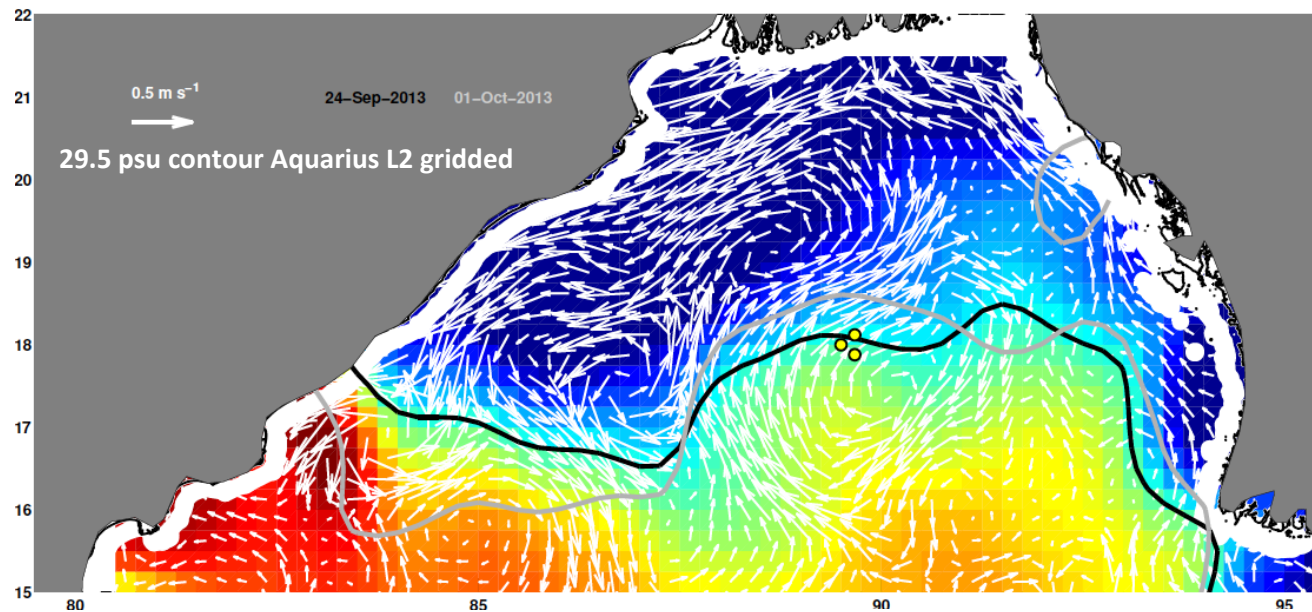


Ganga-Brahmaputra-Meghna water takes a longer path along the western boundary.

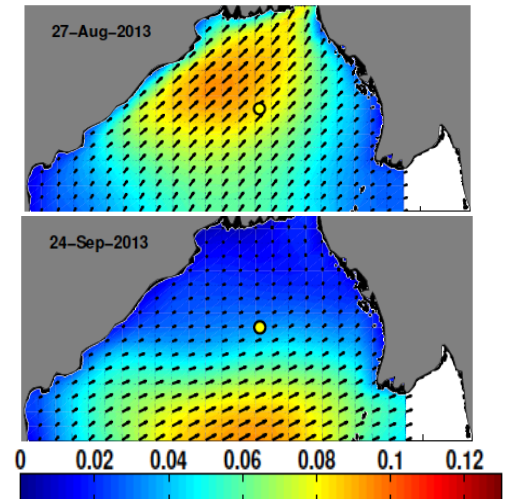
Murty et al., 1992 Gopalakrishna *et.al* 2001 Vinayachandran *et.al* 2002



# Eddy Stirring: Salting episode 2013

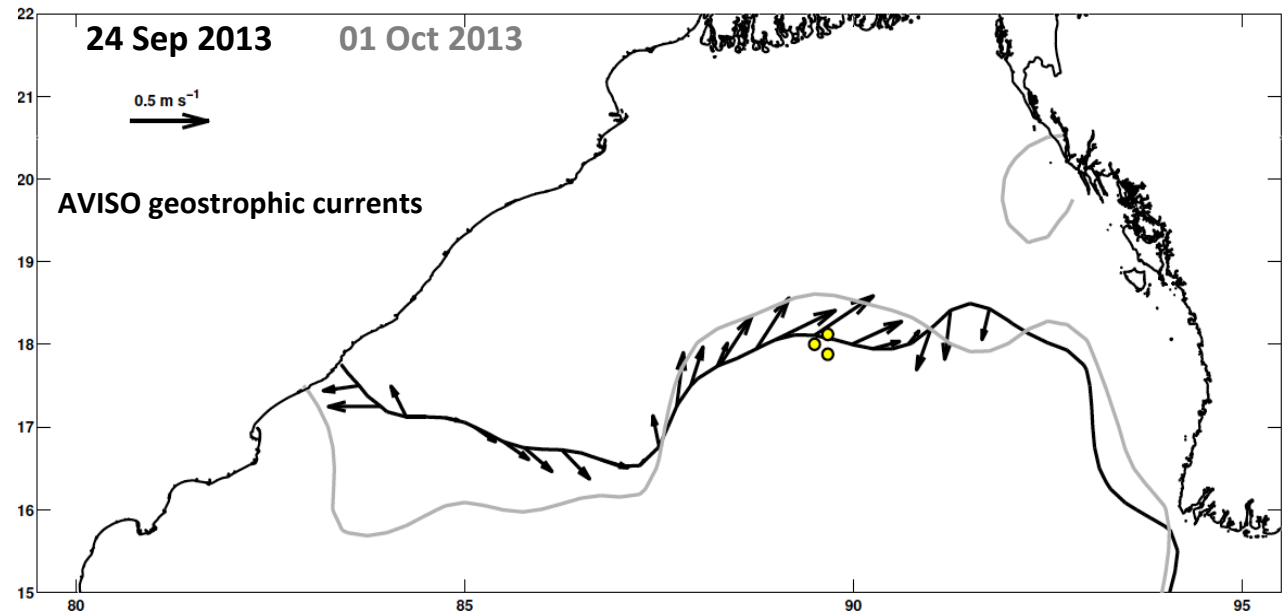


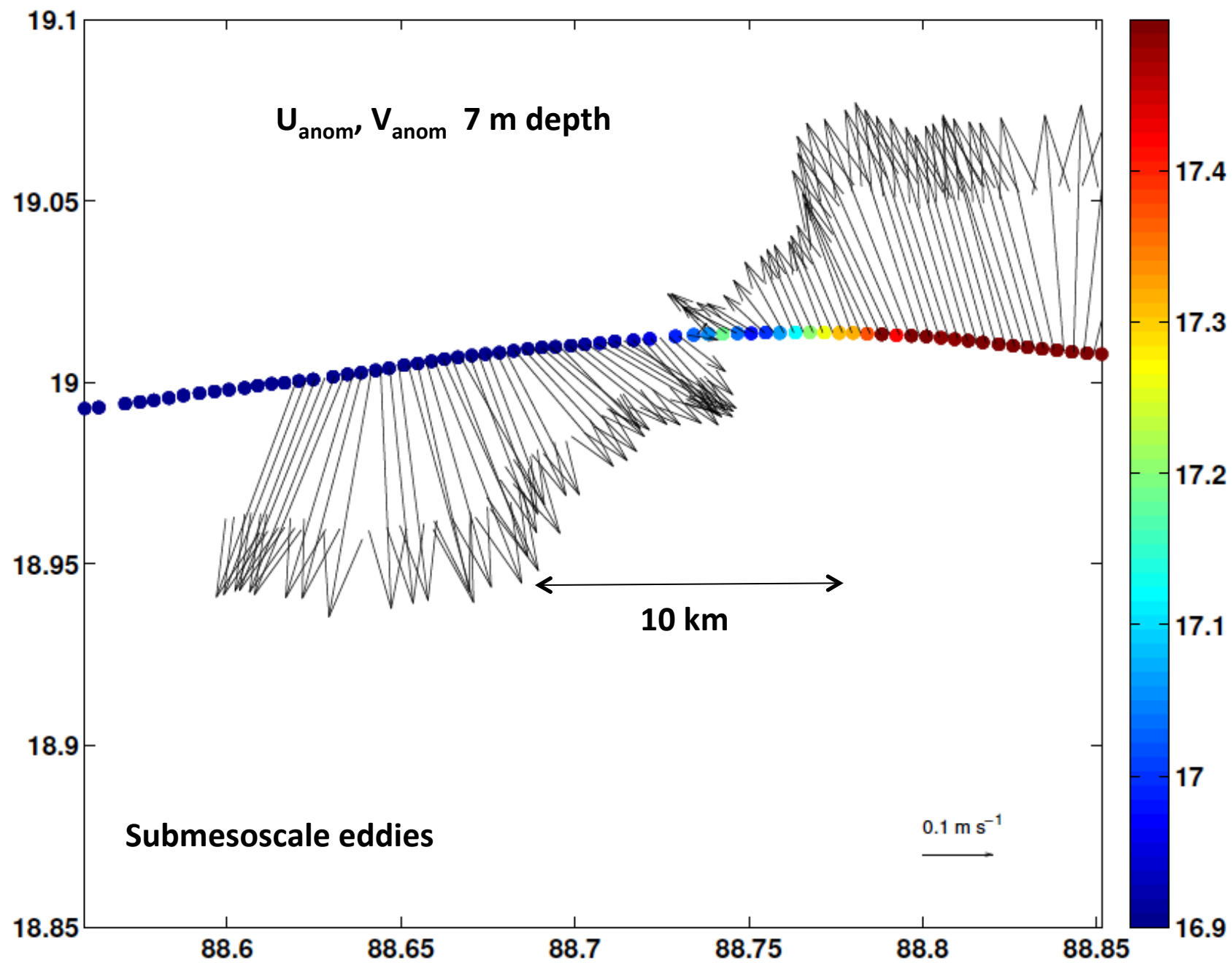
MERRA Wind stress (N/m<sup>2</sup>)



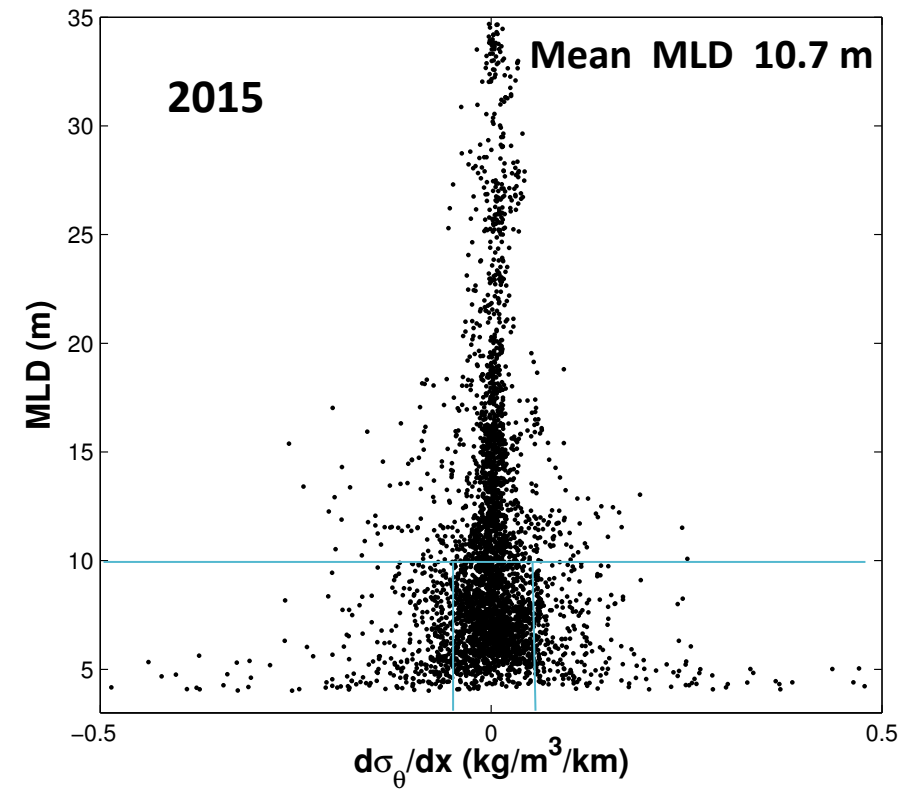
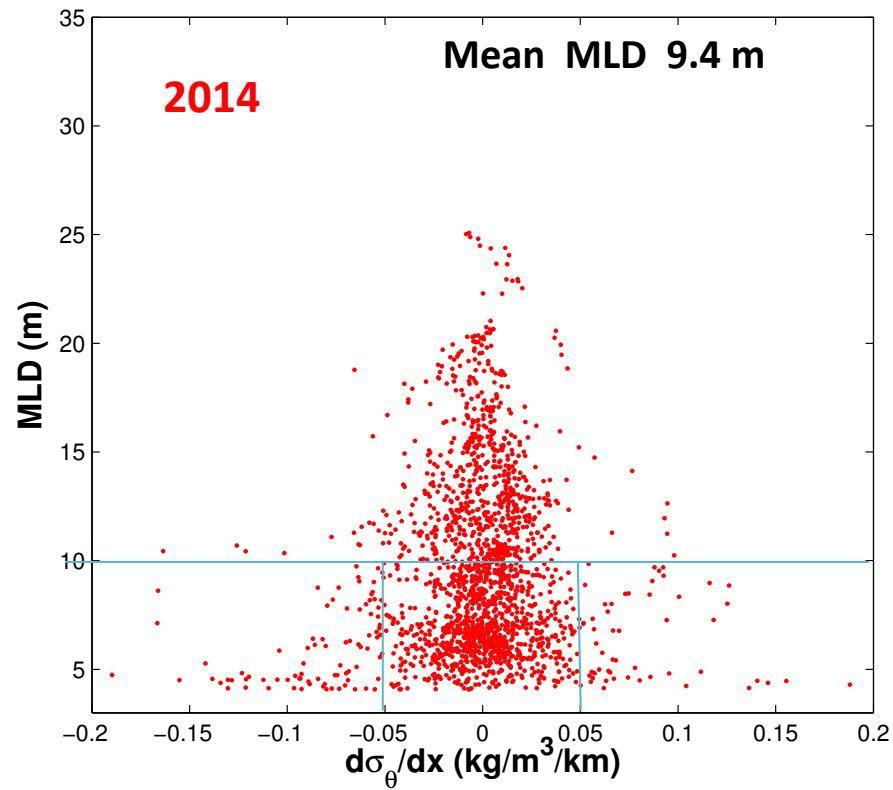
Weak winds 5.2 m/s

Advection by geostrophic currents







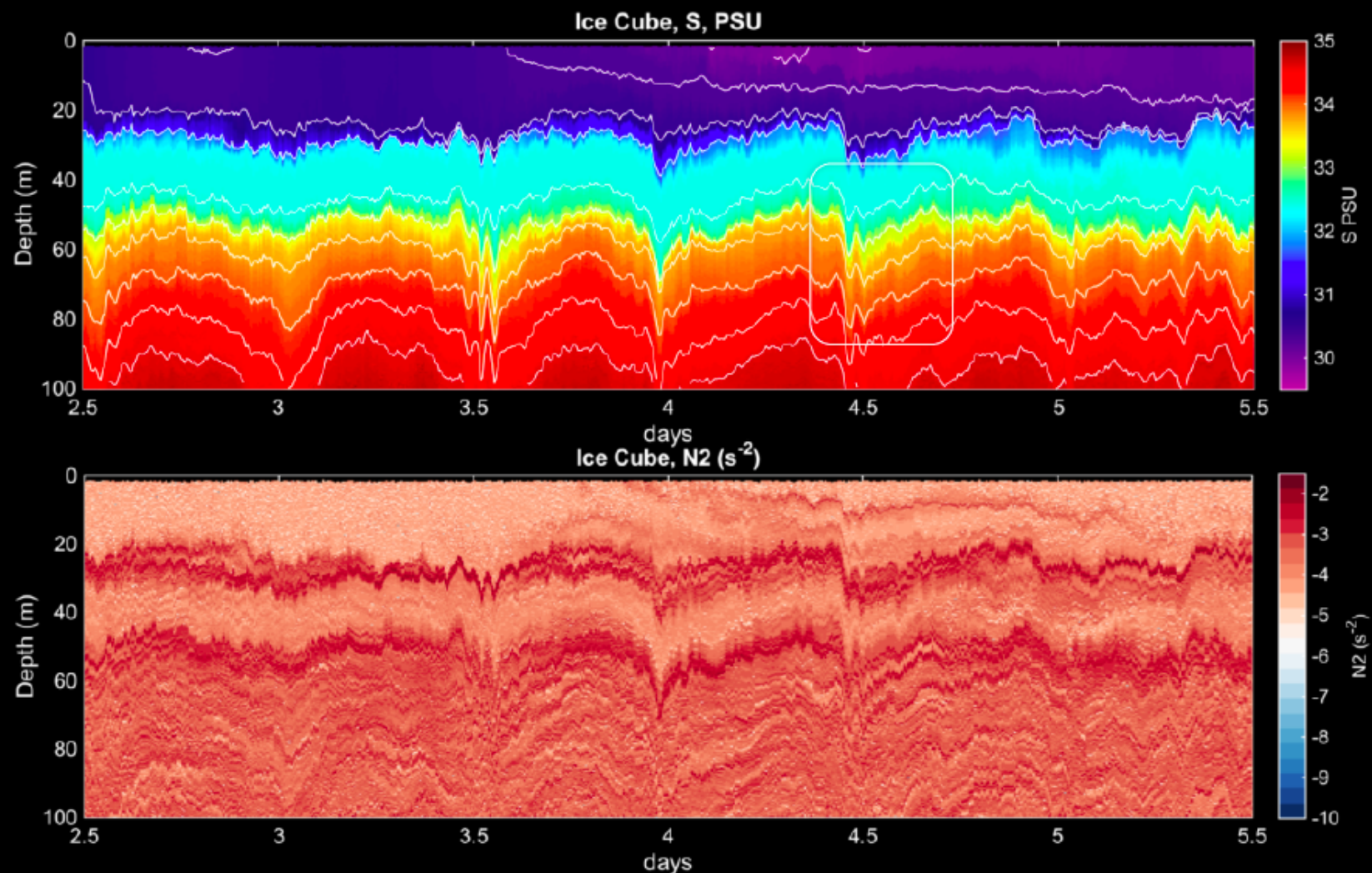


$$\sigma_{\text{mld}} - \sigma_{4\text{m}} = 0.03 \text{ kg/m}^3$$

	$ d\sigma_\theta/dx  > 0.05$		$ d\sigma_\theta/dx  < 0.05$	
MLD Mean	7 m	8 m	10 m	12 m
MLD Stdev	3 m	3 m	4 m	6 m

If  $|d\sigma_\theta/dx| > 0.05 \text{ kg/m}^3 \text{ per km}$ , mixed layer depth is less than 10 m 84 % of the time

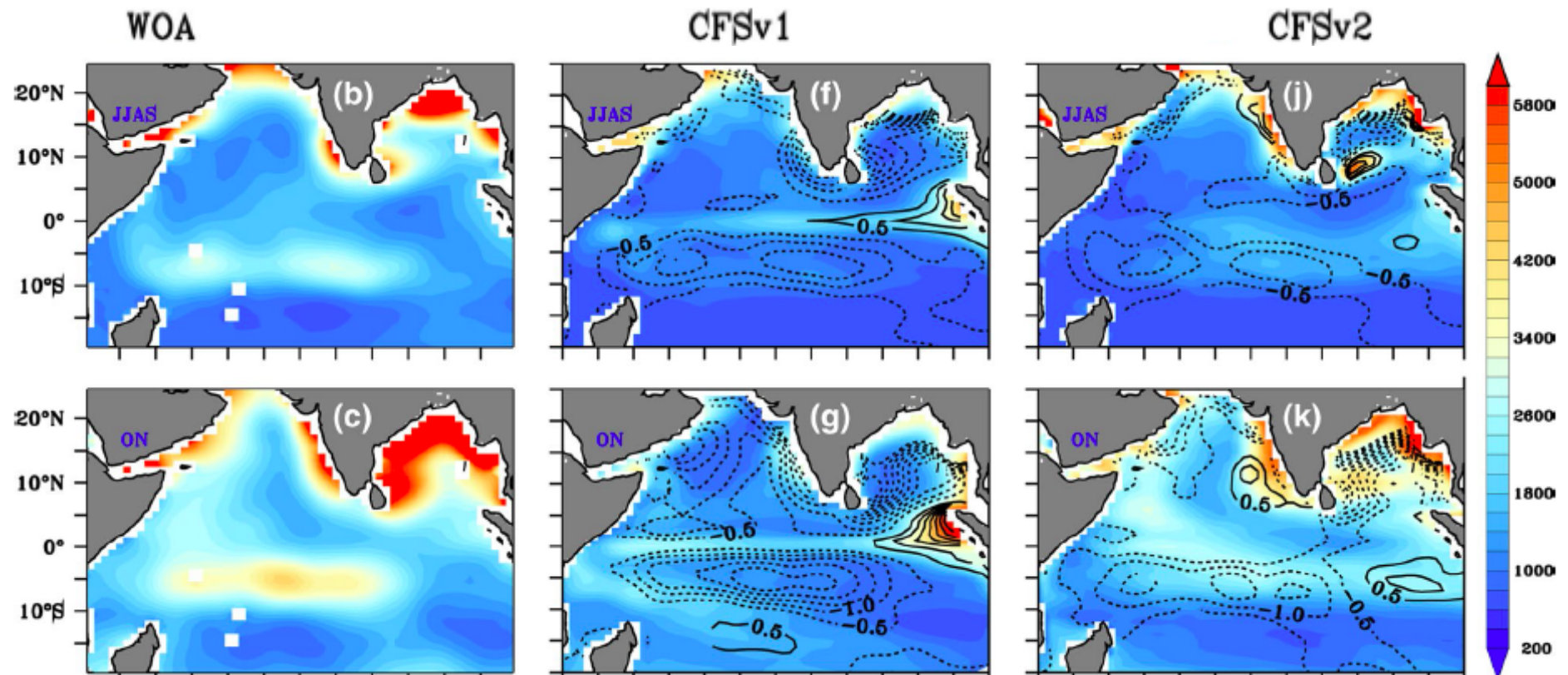
## Wirewalker profiles shows many interesting small scale features



Steep internal tide, perhaps soliton-like waves coincident with surface front by chance. Amazing, persistent small scale structure in main pycnocline (below 60m)

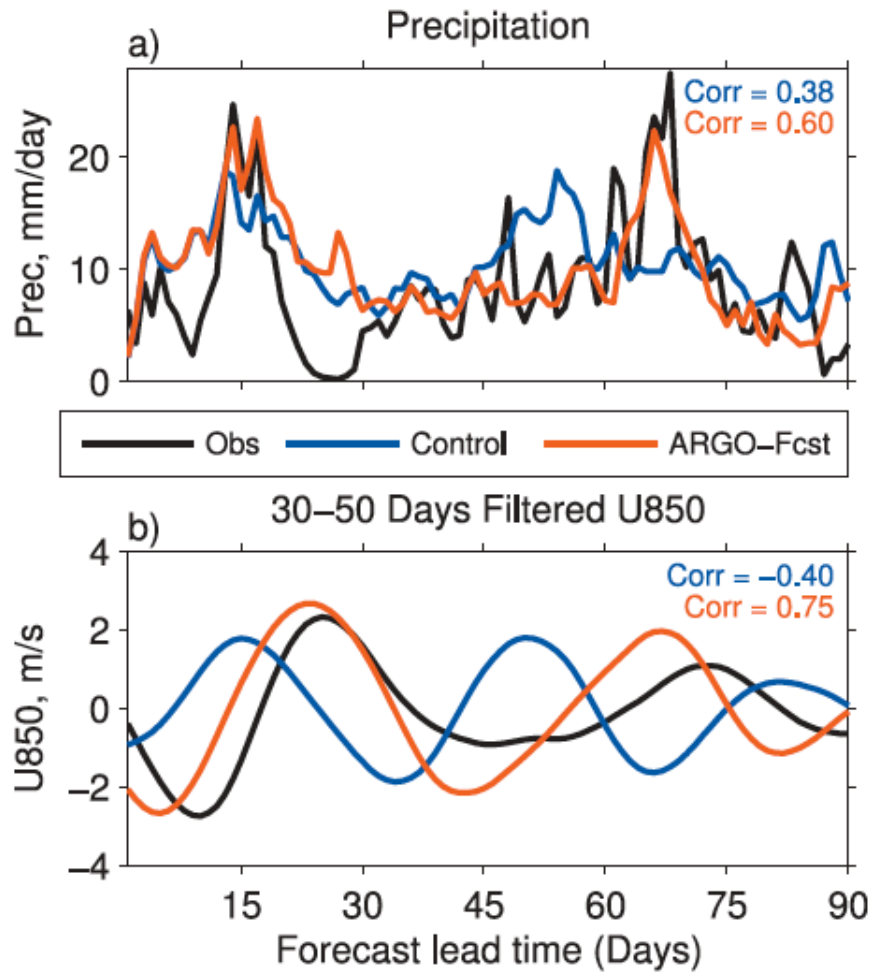
*Courtesy: Drew Lucas*

## Energy required for mixing (ERM) $\text{J}/\text{m}^2$ ; model ERM bias in contours



Parekh *et.al* 2015



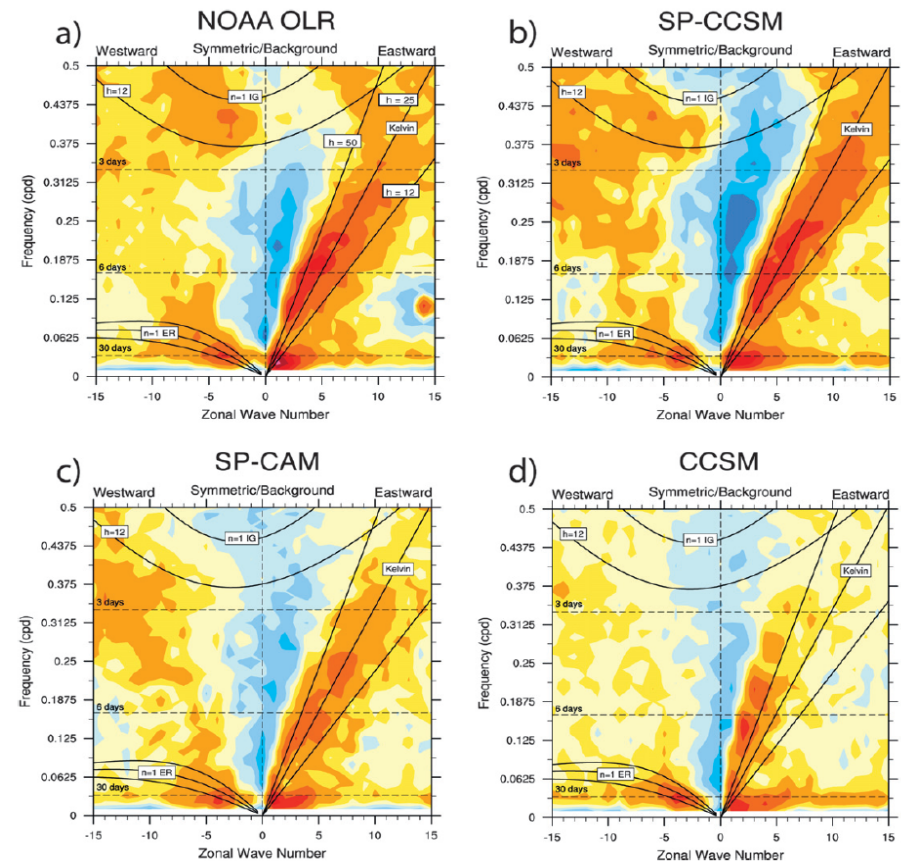


Assimilation of global ARGO subsurface obs to create ocean ICs improves forecast skill

Krishnamurti *et.al* 2007

P and U850 70-90E 15-25N

30 May – 27 Aug 2004



Coupling improves CCEW

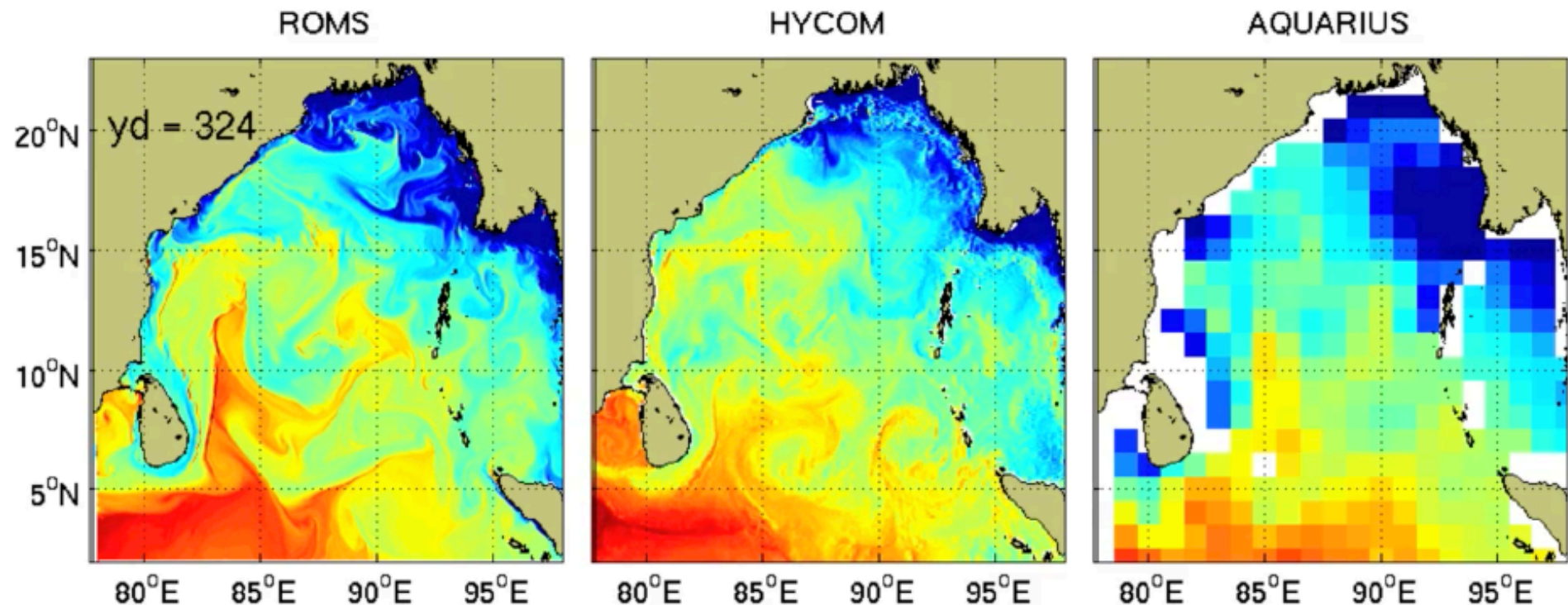
DeMott *et. al* 2011



## BoB variability at many scales is poorly known: Modeling

BoB Eddy rich at  $O(100\text{km})$  scales

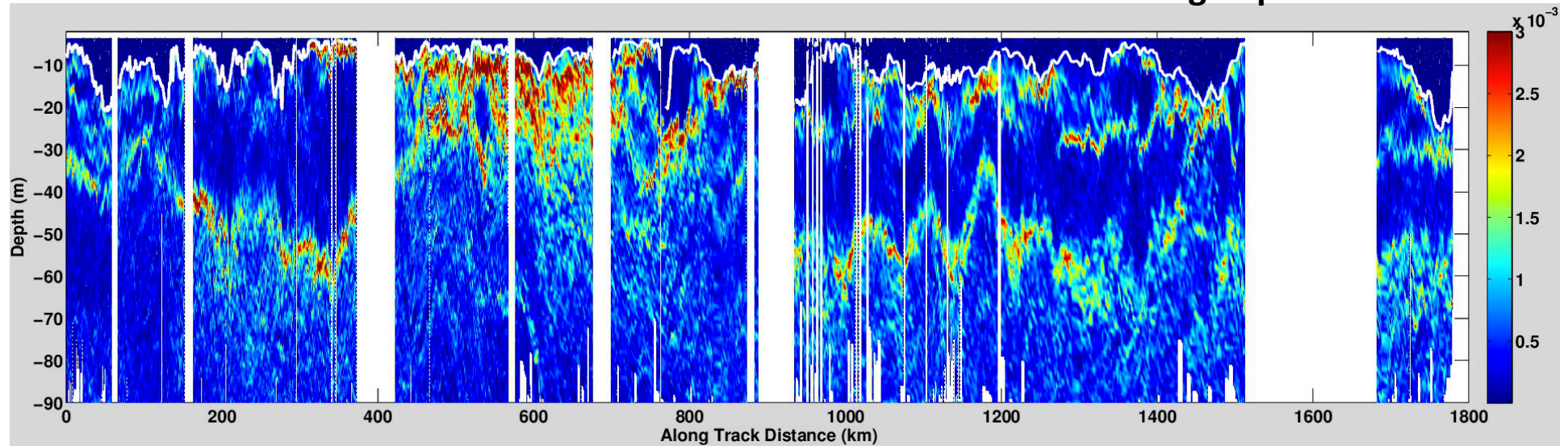
Hints of strong horizontal salinity and density gradients at  $O(1-10\text{km})$  scales



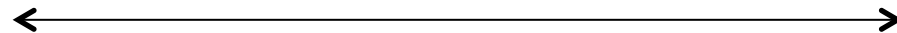
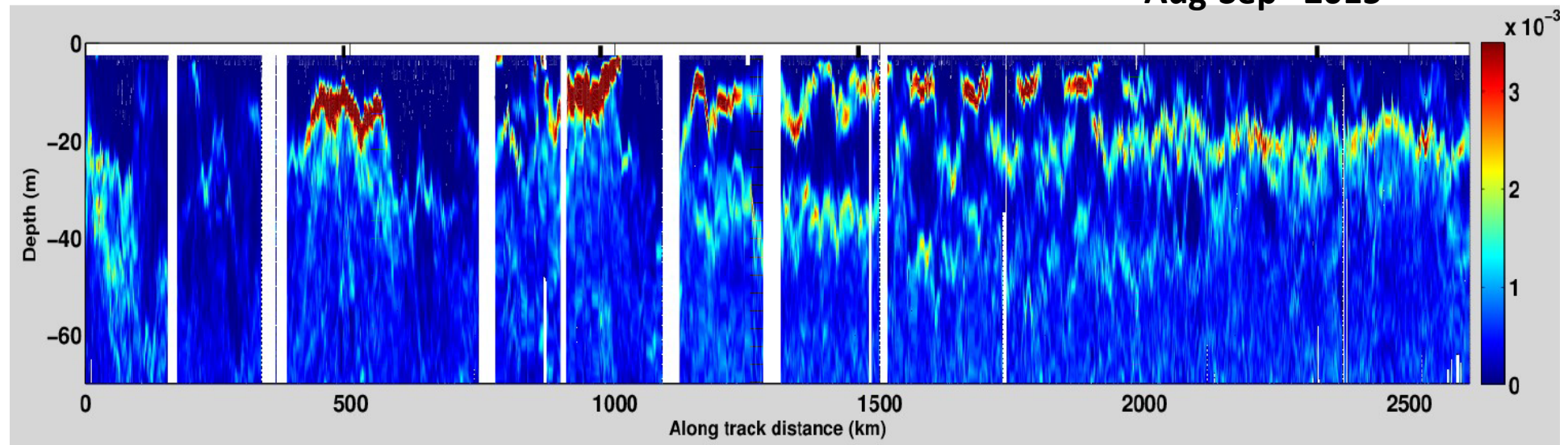
Models and remote imagery often miss stirring & mixing at smaller scales, which have a strong impact on stratification

**Brunt Vaisala Frequency**  $N^2 = -g/\rho (dp/dz)$

**Aug-Sep 2014**

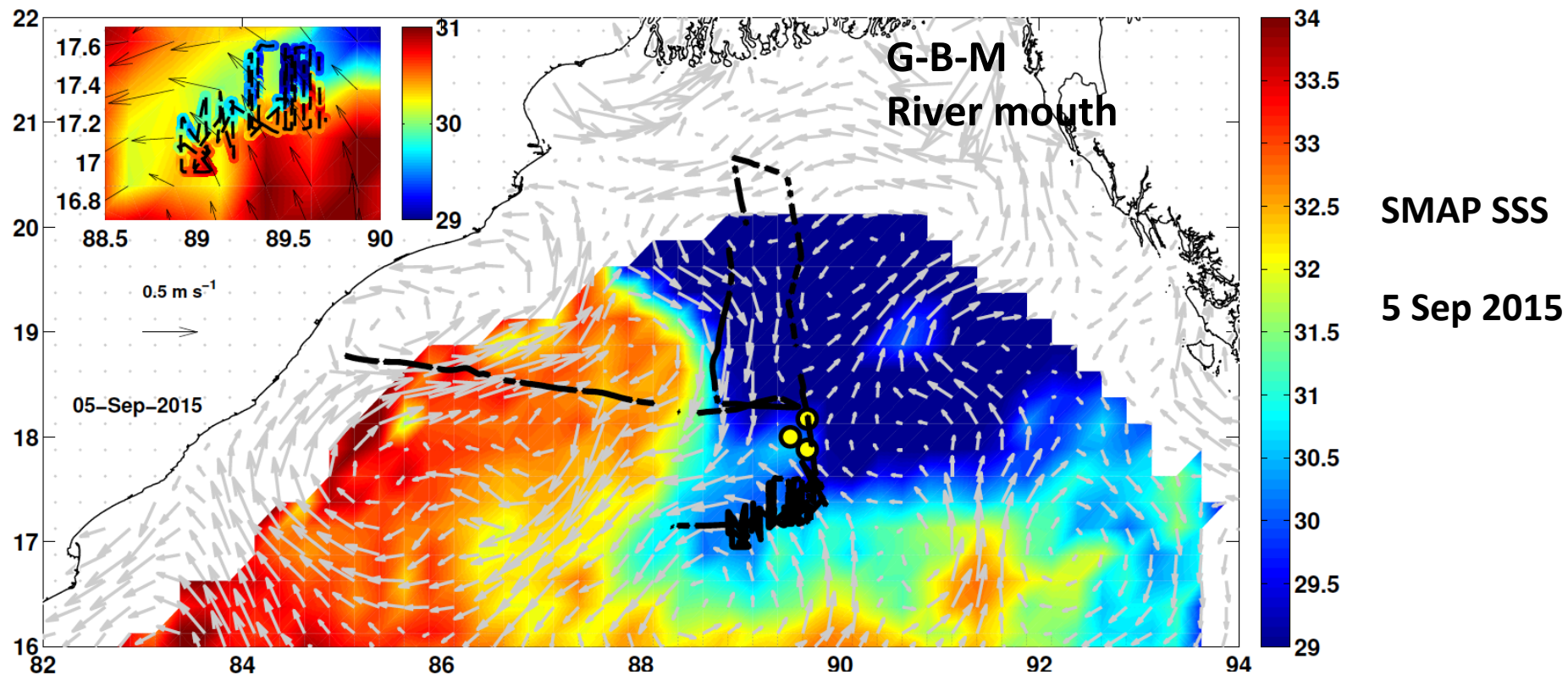


**Aug-Sep 2015**



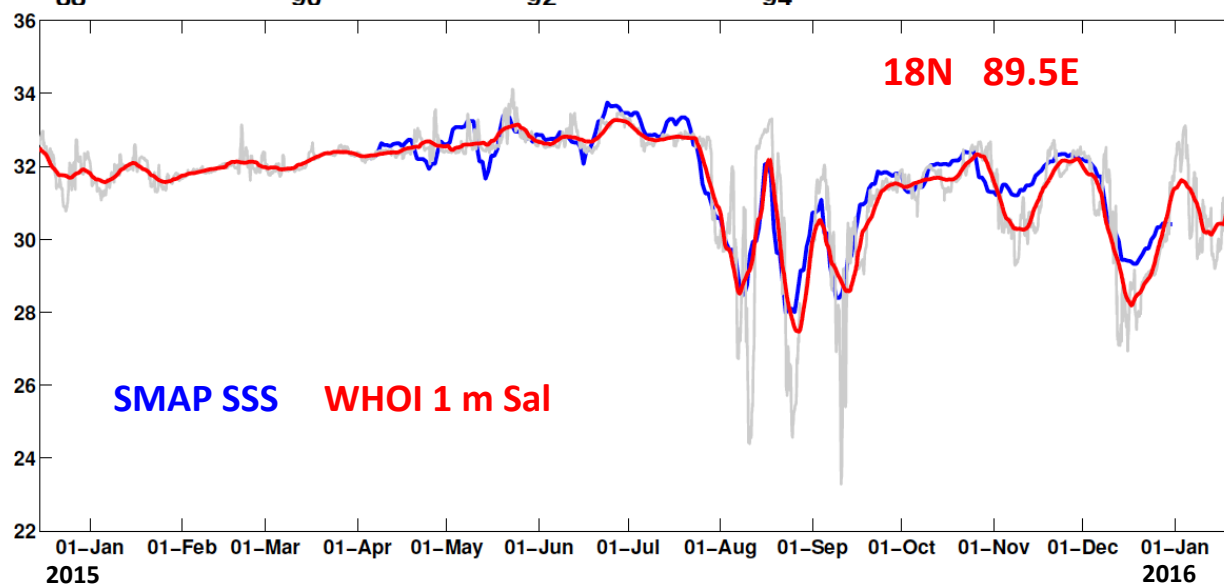
**Two ship experiments**

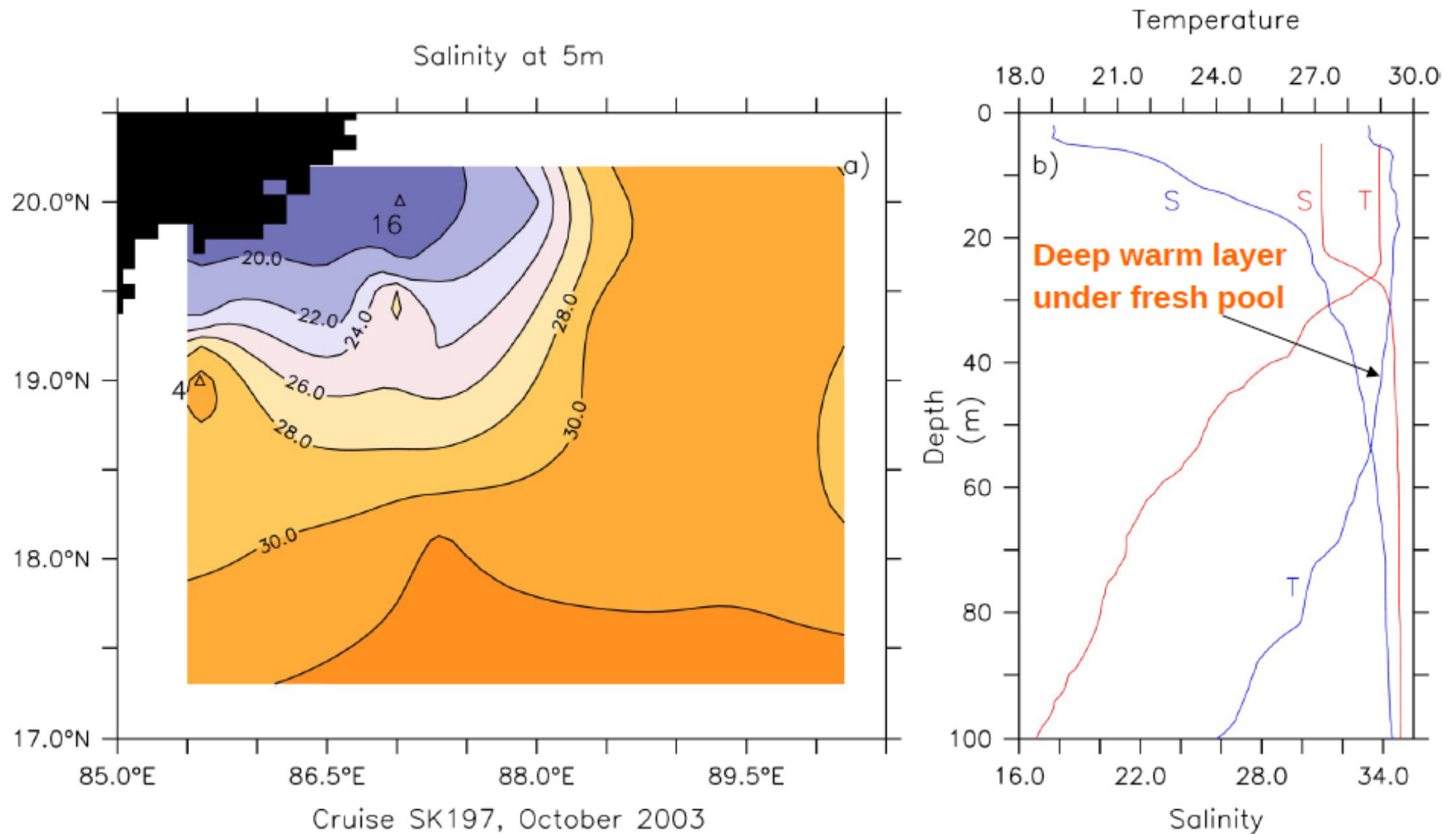




**Soil Moisture Active Passive  
(SMAP) L3 data**

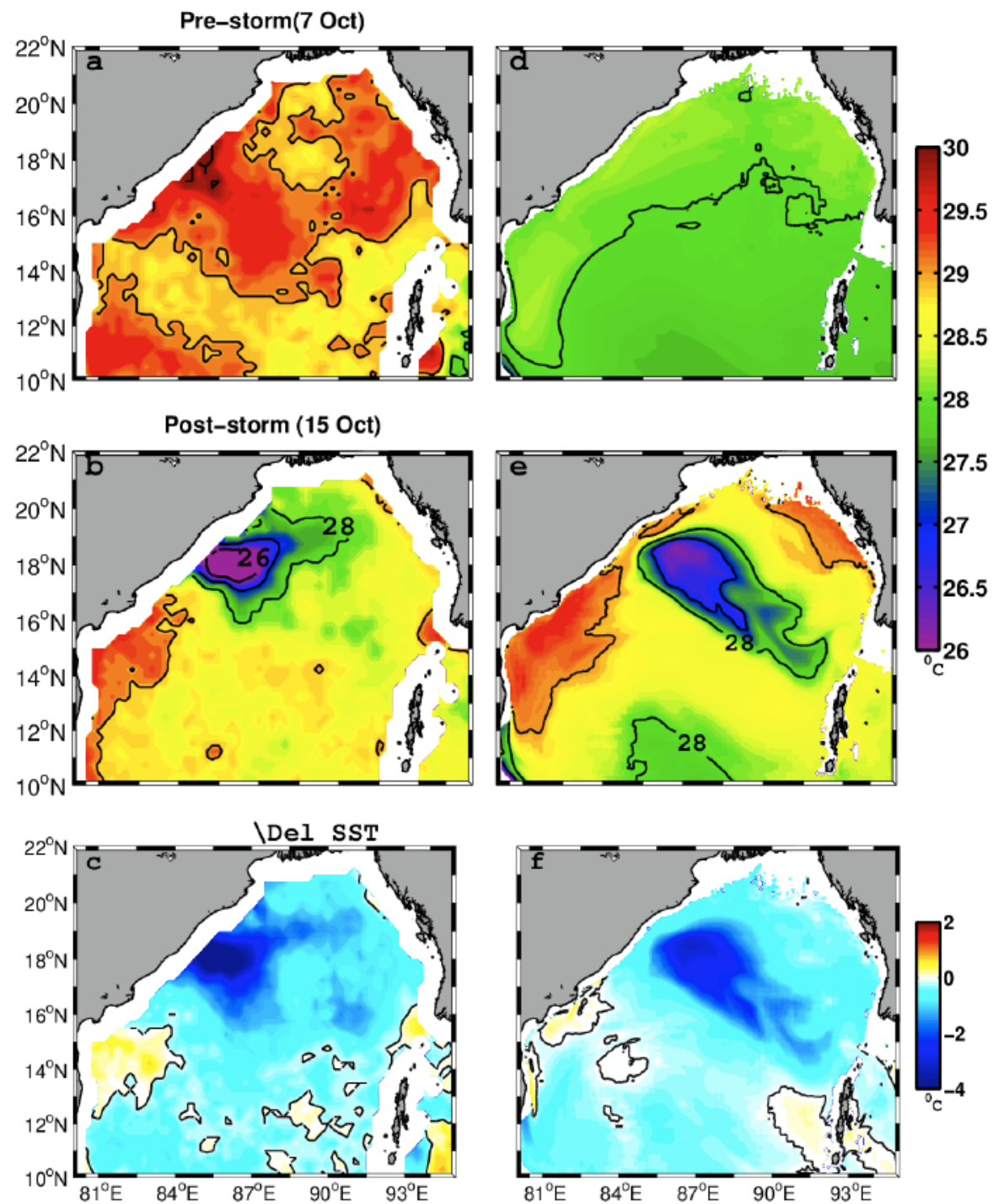
**0.25° x 0.25° resolution; 8 day  
running averages**





**Deep warm layer under fresh pool due to penetrative sunlight and surface cooling**







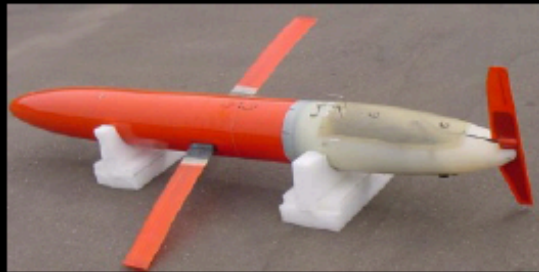
Shipboard training and research on joint cruises: Nov-Dec 2013, June 2014, August 2014, November 2014, August-September 2015



Bay of Bengal Upper Ocean Physics Workshop IISc Bengaluru July 9-21, 2014  
Also, Marine Mammal Workshop NCBS India by ASIRI PIs

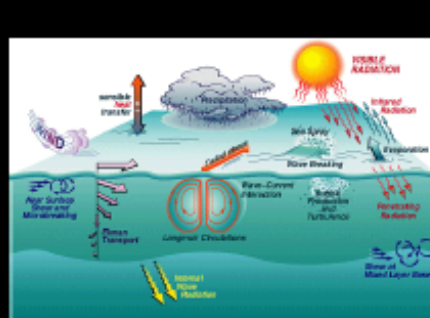
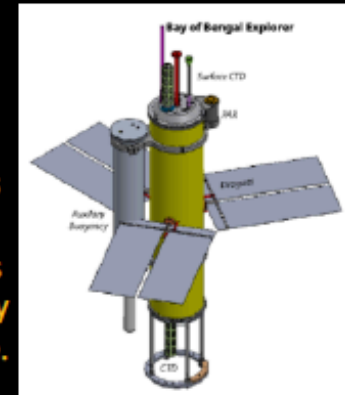


# Observational Tools



Long term endurance gliders Spray (left) and Seaglider (right) measure subsurface temperature and salinity structure.

Lagrangian float follows density or pressure levels while taking measurements and is remotely programmable.



Air-sea flux mooring measures air-sea exchange of heat, momentum, and freshwater and temperature, salinity, and velocity in the ocean.

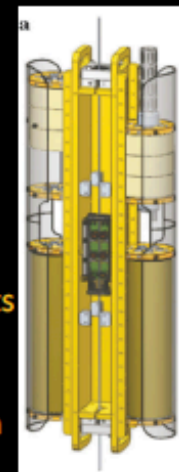


Turbulent motions are being measured from moorings, CTD rosettes, and autonomous profilers using chipods.

Slocum turbulence glider measures microstructure shear and mixing.



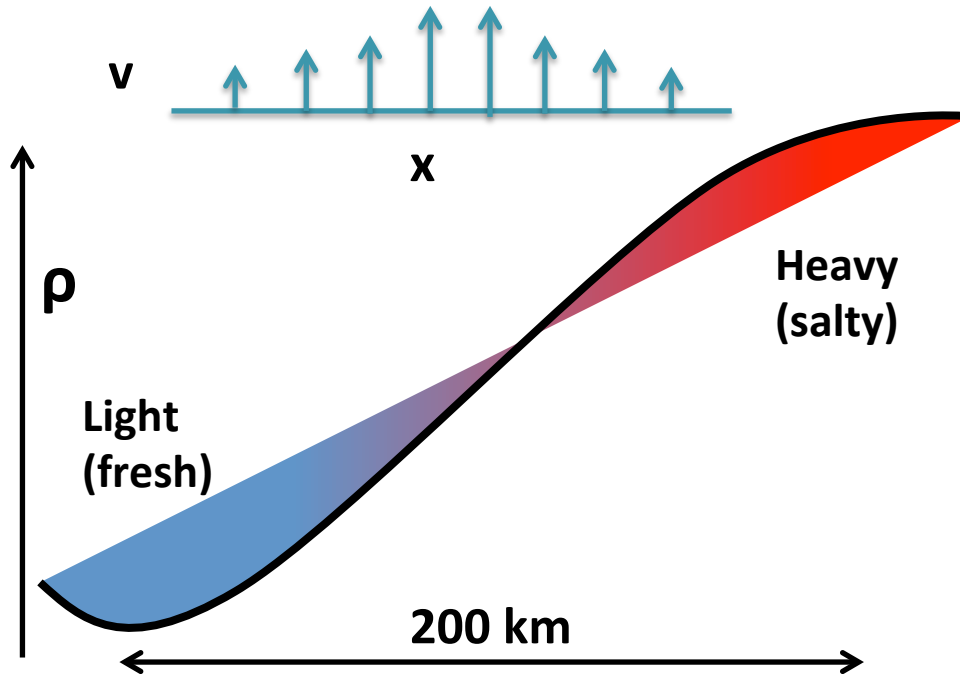
Surface wave-driven wirewalker measures temperature, salinity, currents and optical variables at cm scale resolution



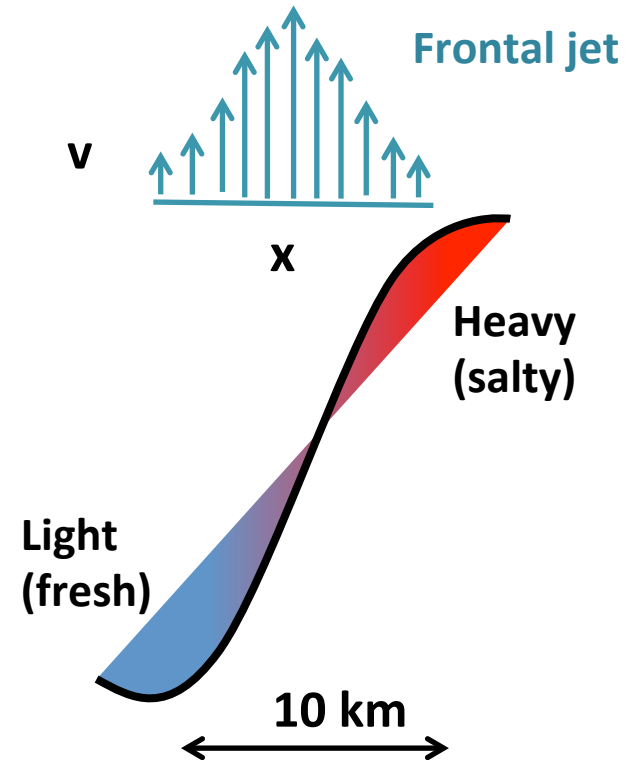
Lagrangian drifters measure near-surface ocean currents, atmospheric pressure, sea surface temperature and sea surface salinity

# Lateral scales

**Mesoscale** Order 100 km  
 $Ro \ll 1$  weak lateral shear



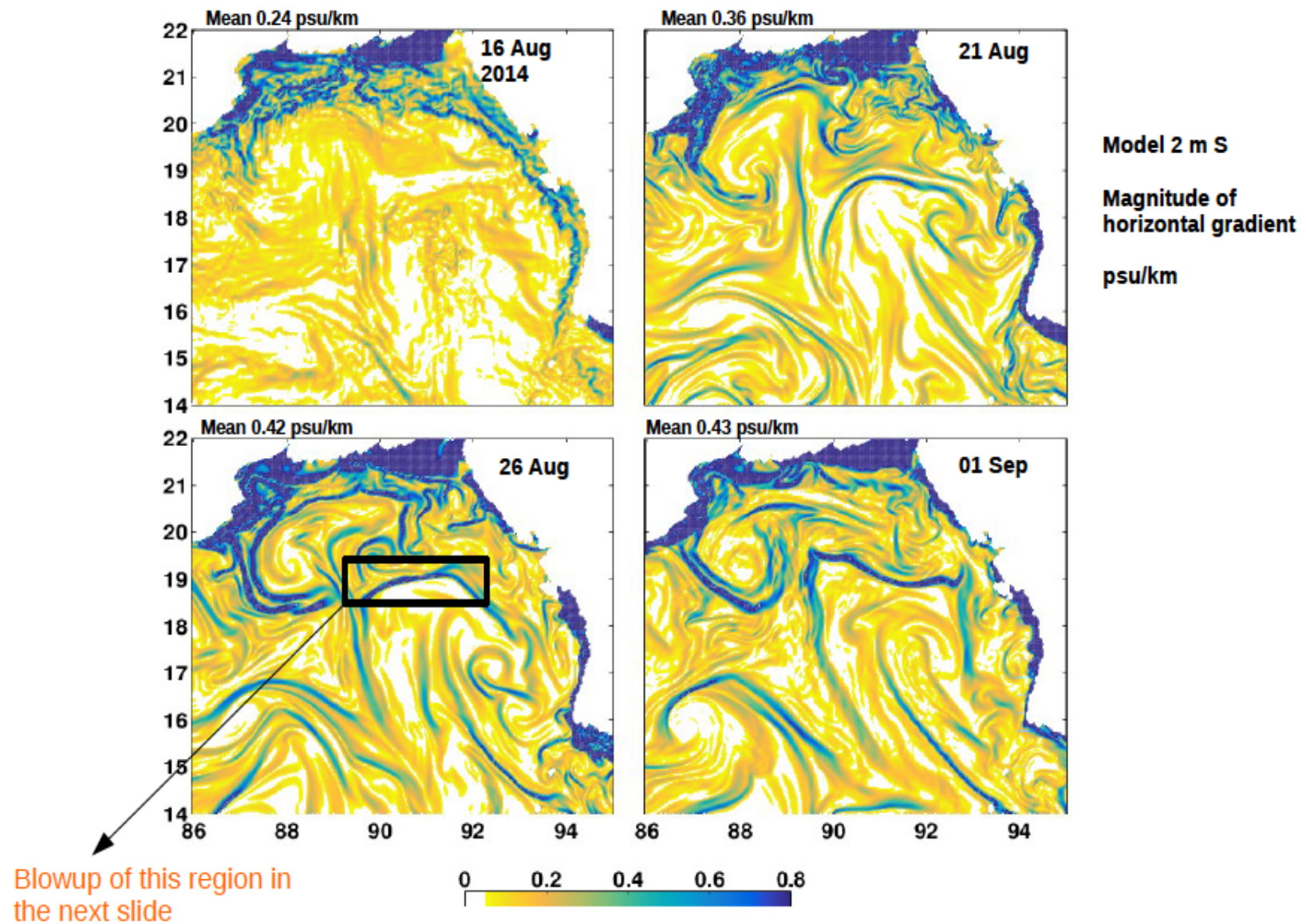
**Sub-mesoscale** Order 1-10 km  
 $Ro \sim O(1)$  strong lateral shear



**Slumping of sub-mesoscale fronts creates shallow stratification,  $T \sim O(1)$  day**

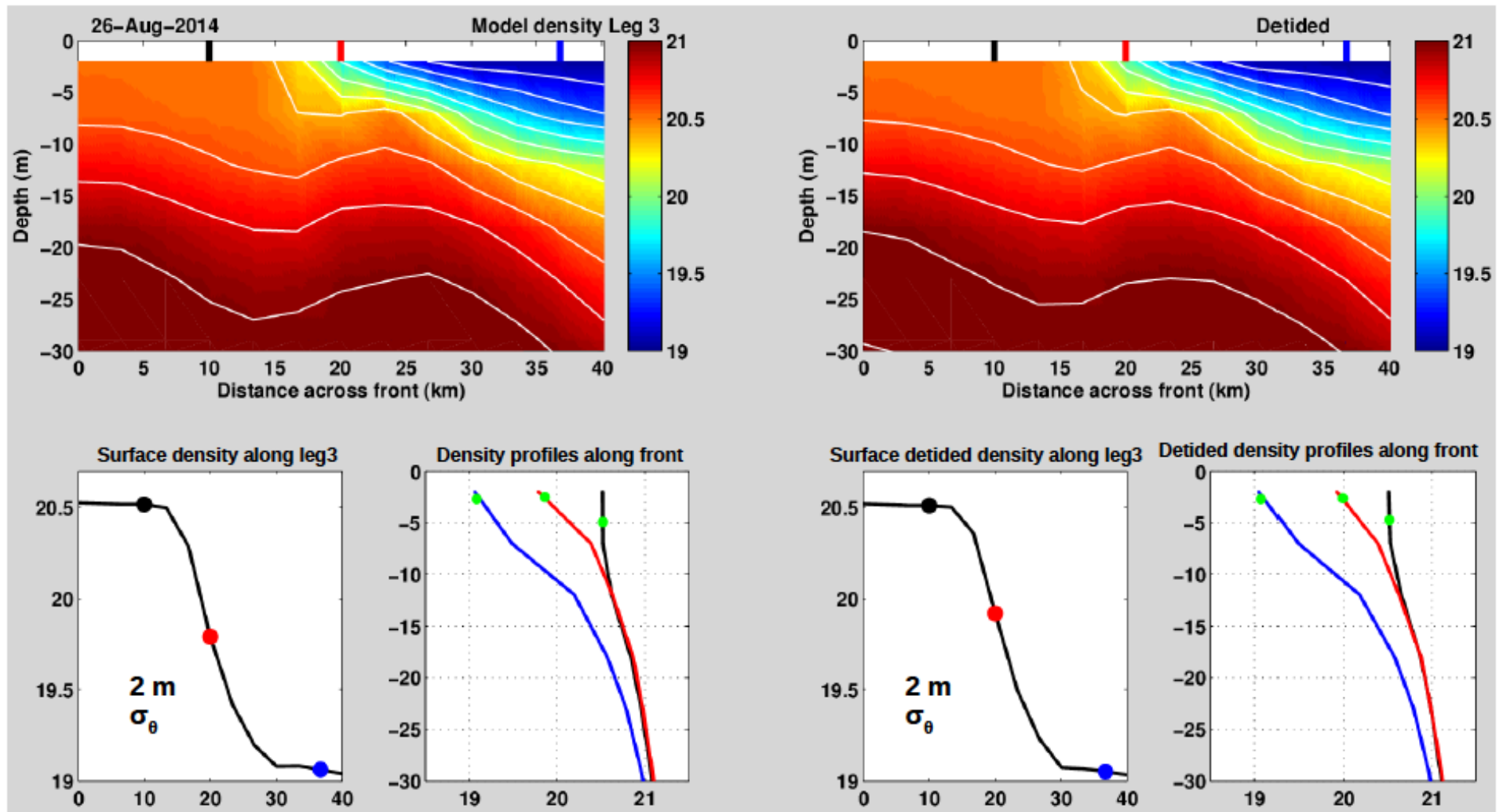
*Tandon Garrett 1995 Mahadevan Tandon 2006  
Thomas 2007 D'Asaro 2013*





Lateral salinity gradients sharpened by mesoscale flow (see slide 1)

The domain average gradient magnitude increases 2 times from  $t=0$  to  $t=15$  days

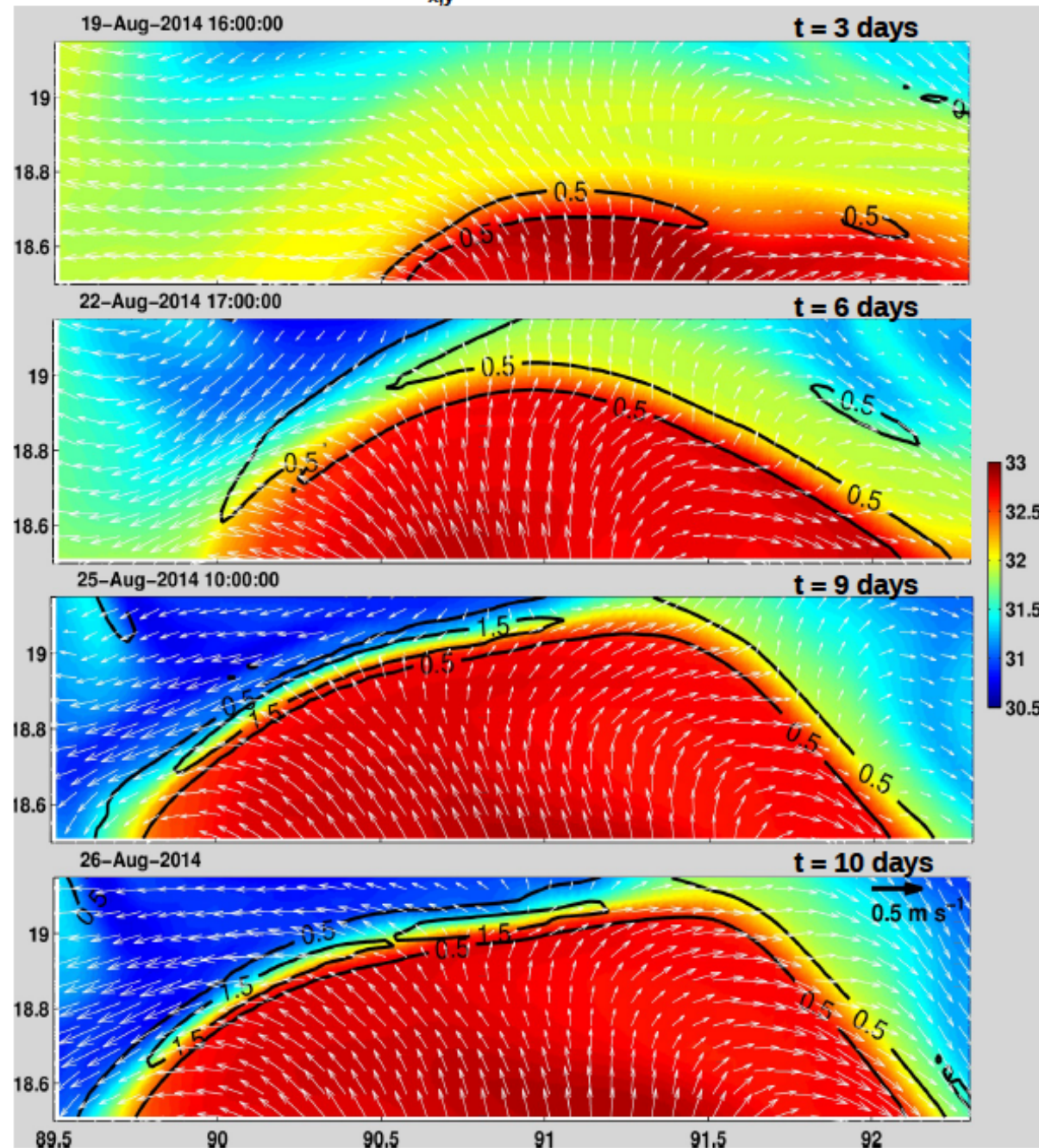


The green dots on the density profiles indicate the mixed layer depth (m), calculated using the criteria  $\sigma_{\text{mld}} - \sigma_{\text{surface}} = 0.3 \text{ kg/m}^3$

The mixed layer is shallow on the fresher side and under the front than on the saltier side.

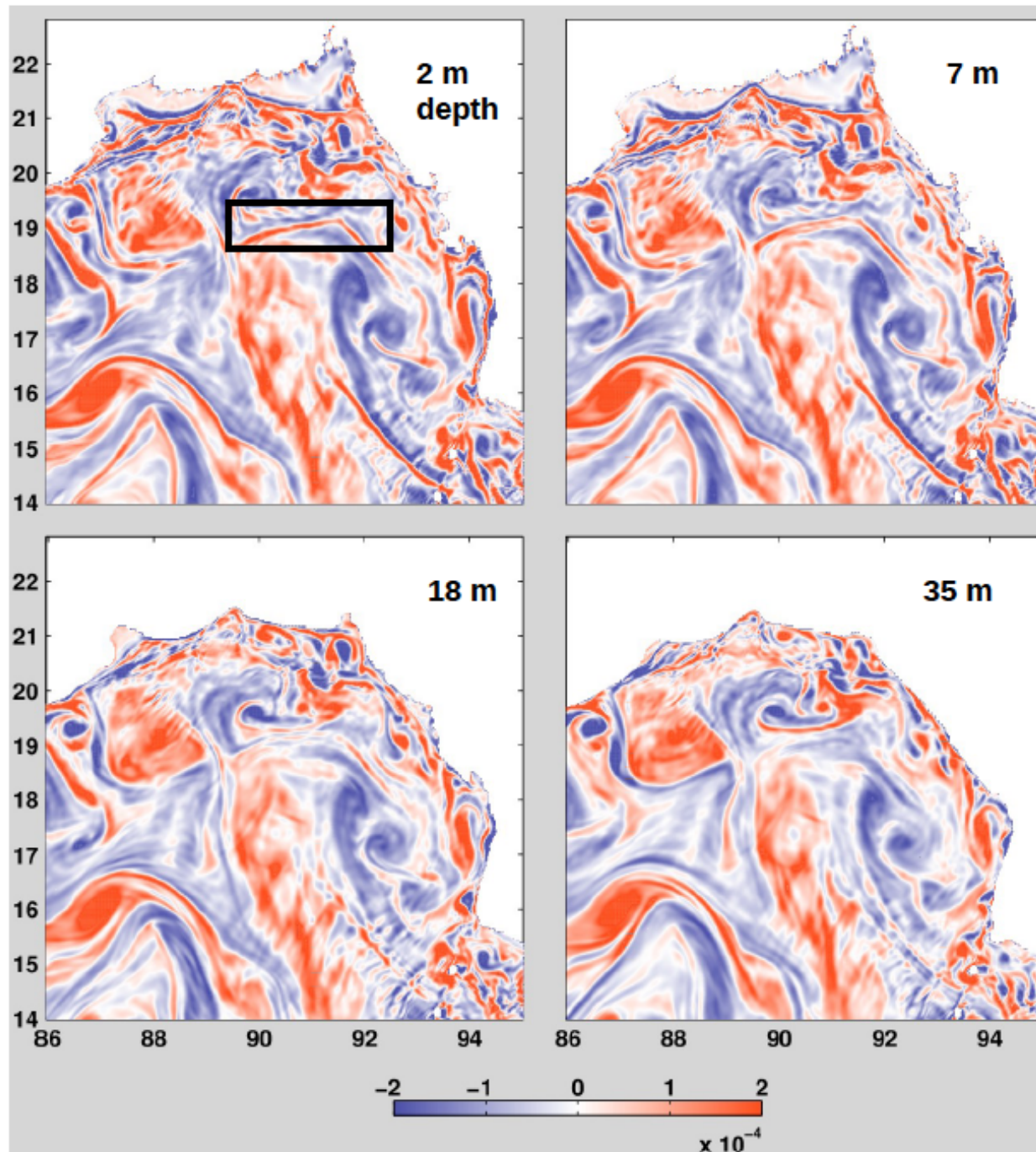


2 m salinity (psu; color);  $|\text{grad}S_{x,y}|$  (psu/km; black contour); current vectors (detided)



The front is formed due to confluence by mesoscale eddies.

By t=10 days, the magnitude of horizontal salinity gradient across the front strengthens (2 psu/km )



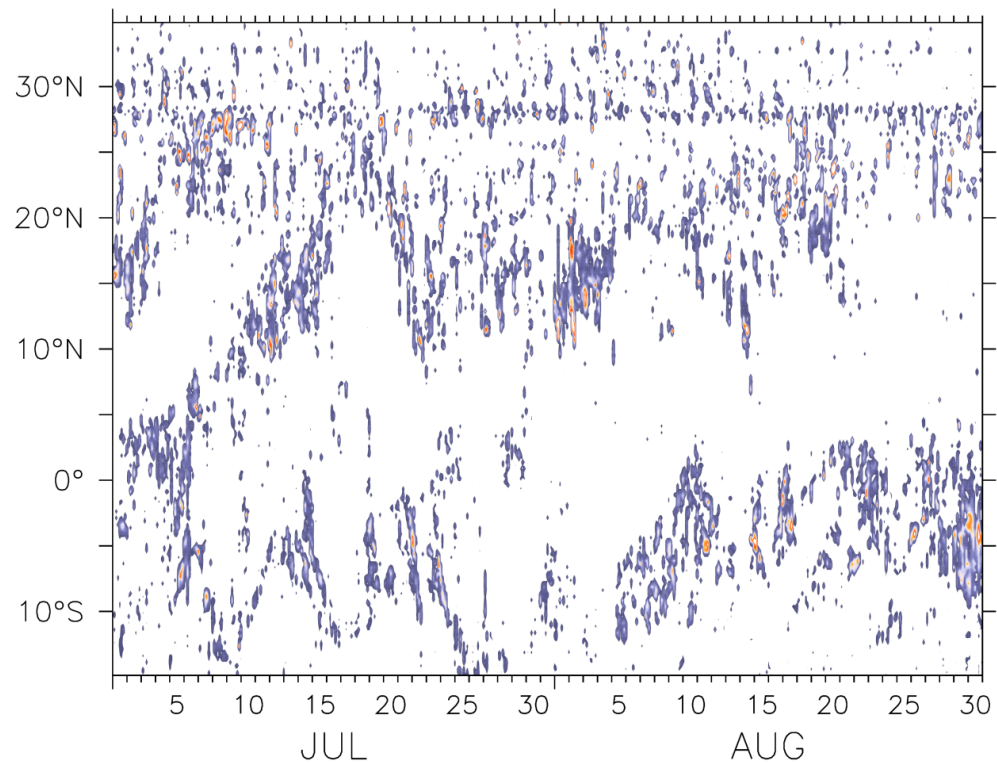
The vorticity at different depths suggest that the submesoscale (order 10 km) frontal regions associated with high positive vorticity are confined to the upper 10 m.



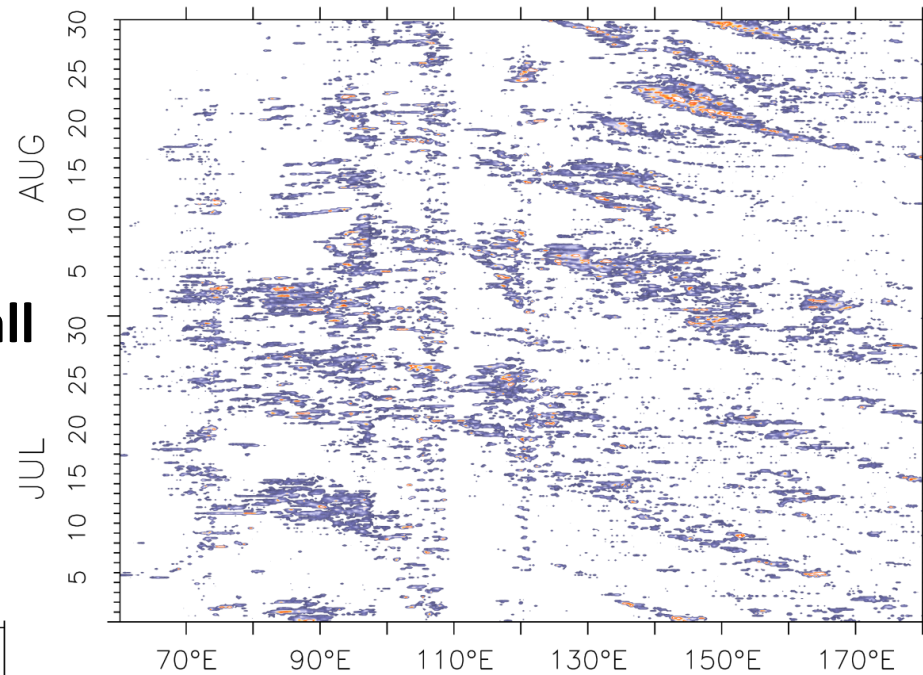
# Jul-Aug 2004 TRMM 3-hourly rainfall

15 N

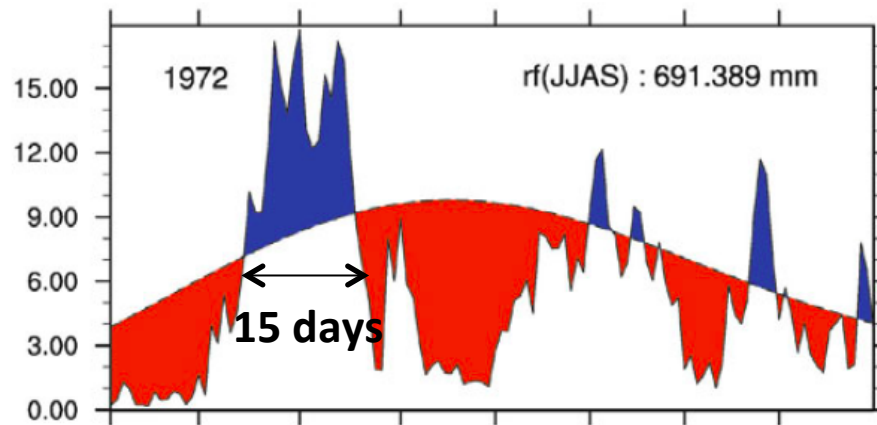
3 Hourly Rainfall at 85.9E ,2004



3 Hourly Rainfall at 14.9N ,2004

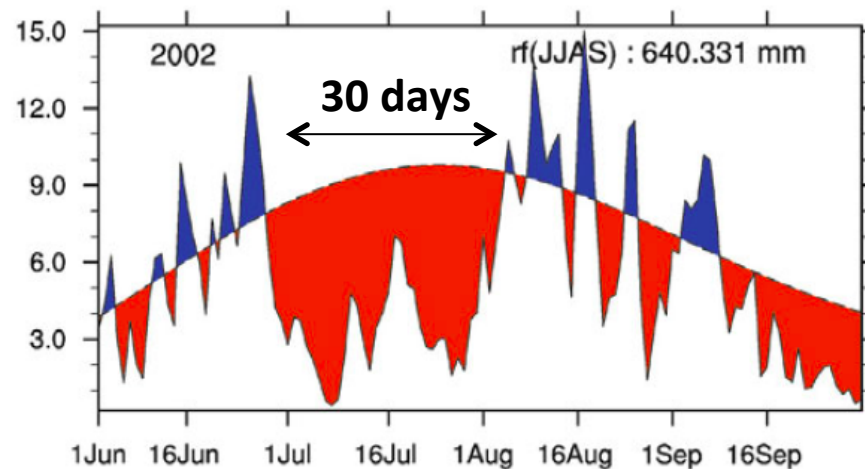
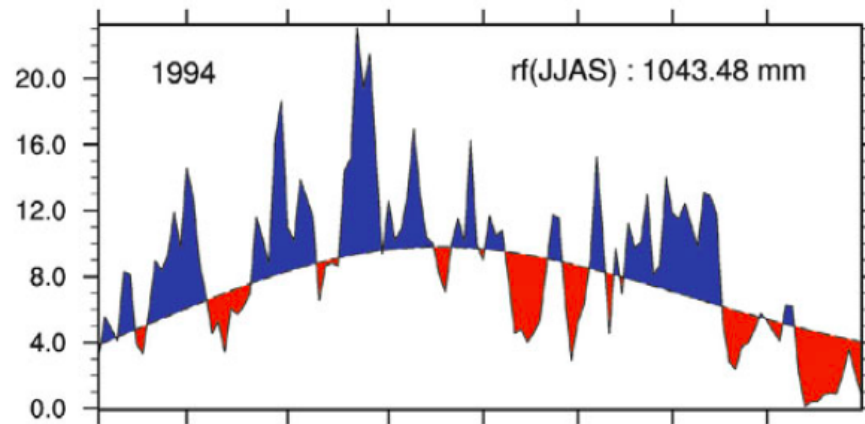


86 E



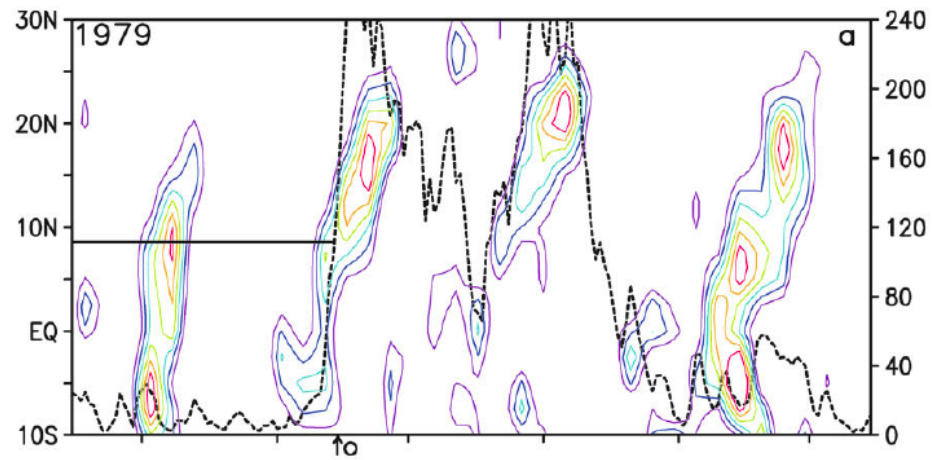
**Daily gauge rainfall (mm/day)  
averaged over Central India**

**72.5E–85.5E 10.5N–25.5N**

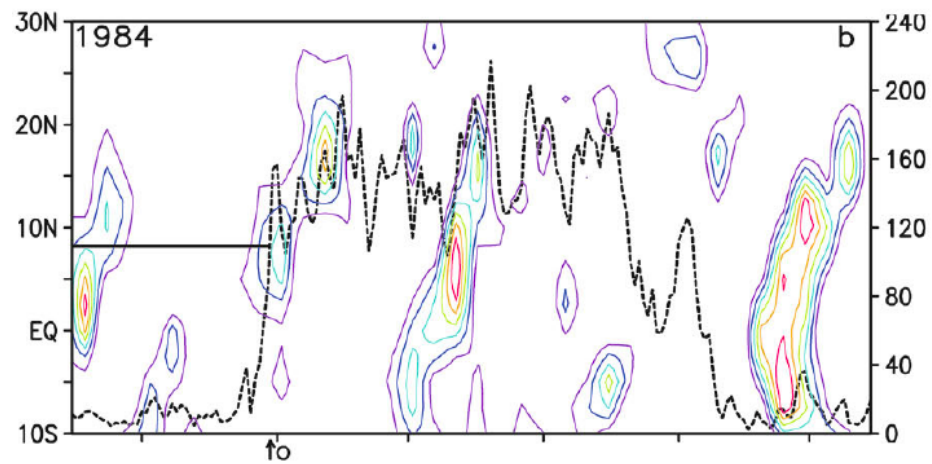


**“Active-Break” cycle of monsoon**

***Goswami 2012***

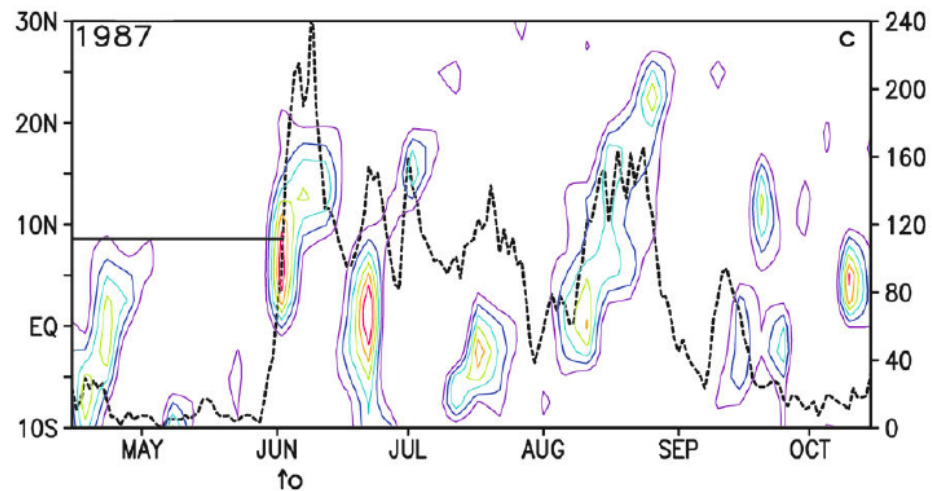


**CMAP Rainfall 70E-90E**



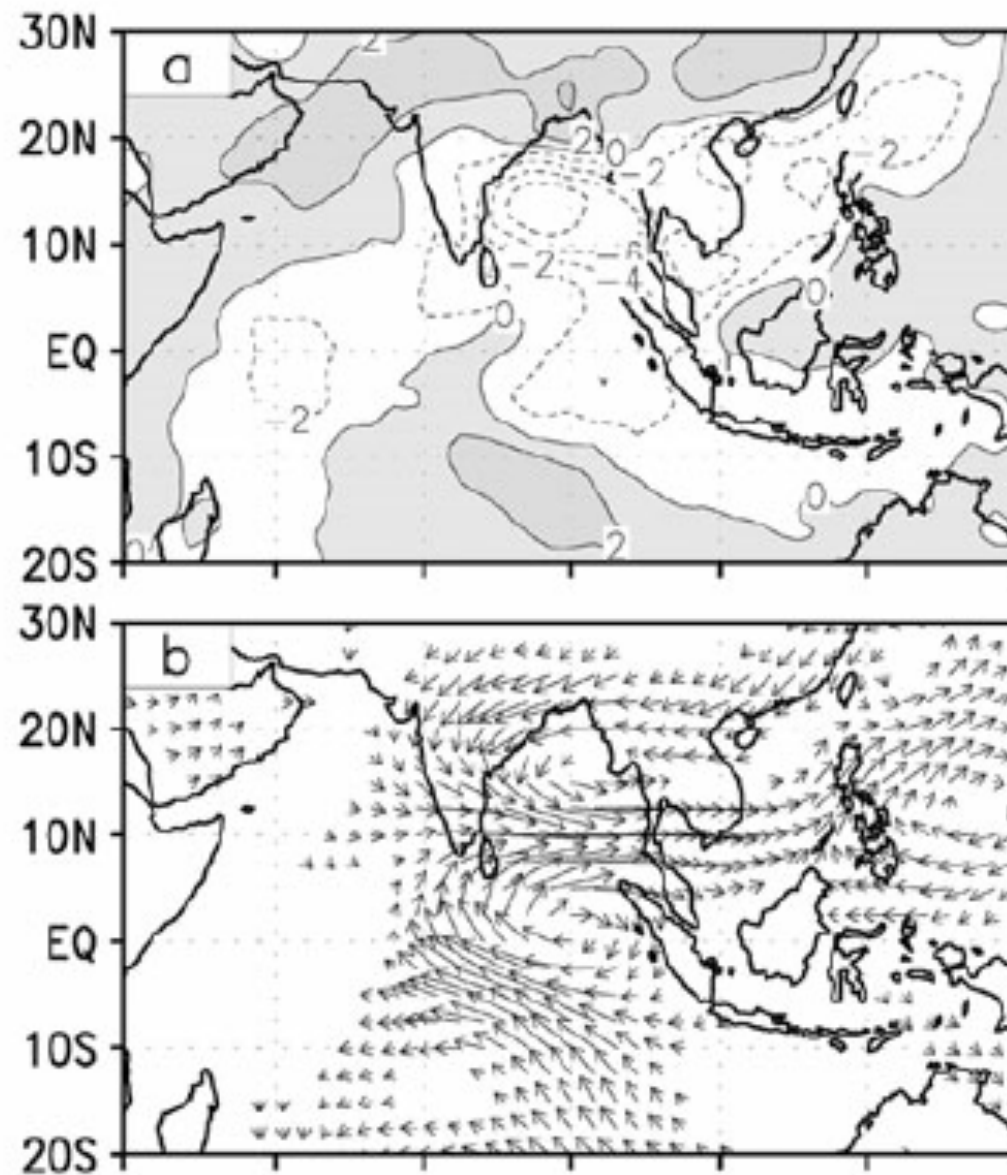
**KE 850 hPa winds 55E-65E 5N-15N**

**30-50 day mode**



***Goswami 2012***

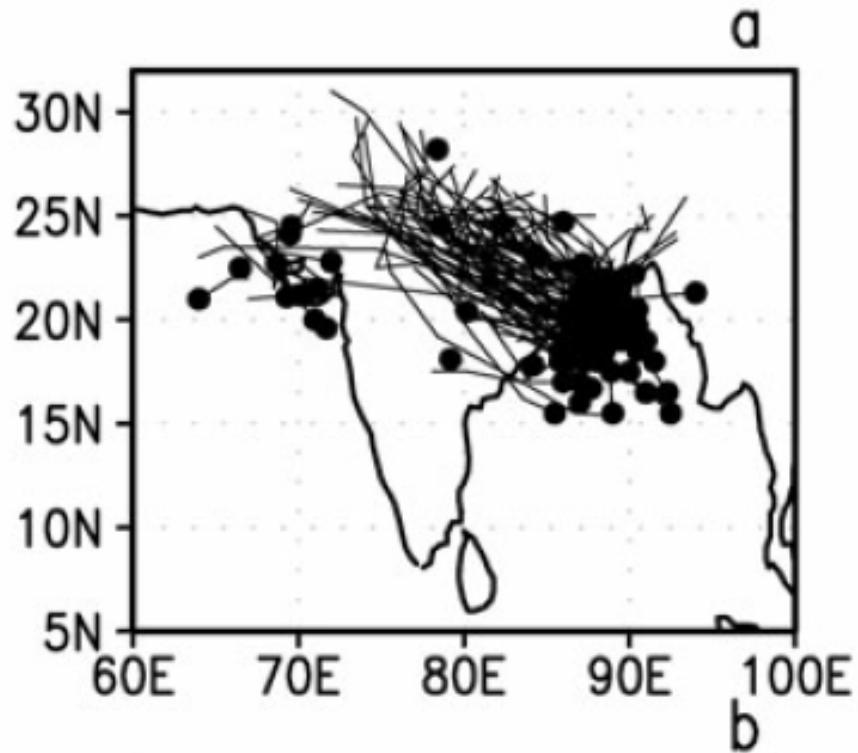




**OLR 850 hPa winds**

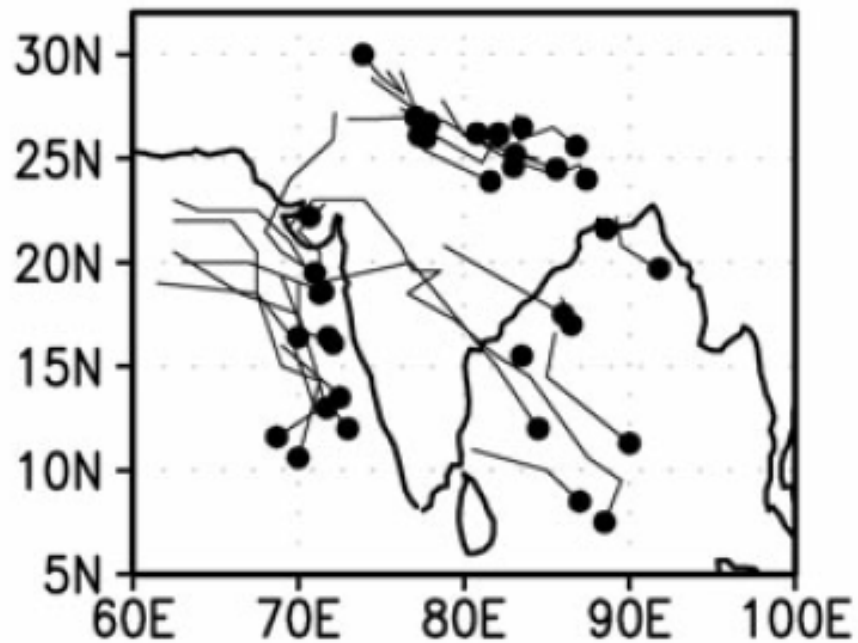
**10-20 day mode**

***Goswami 2012***



Tracks of Monsoon low-pressure systems/depressions 1954-1983

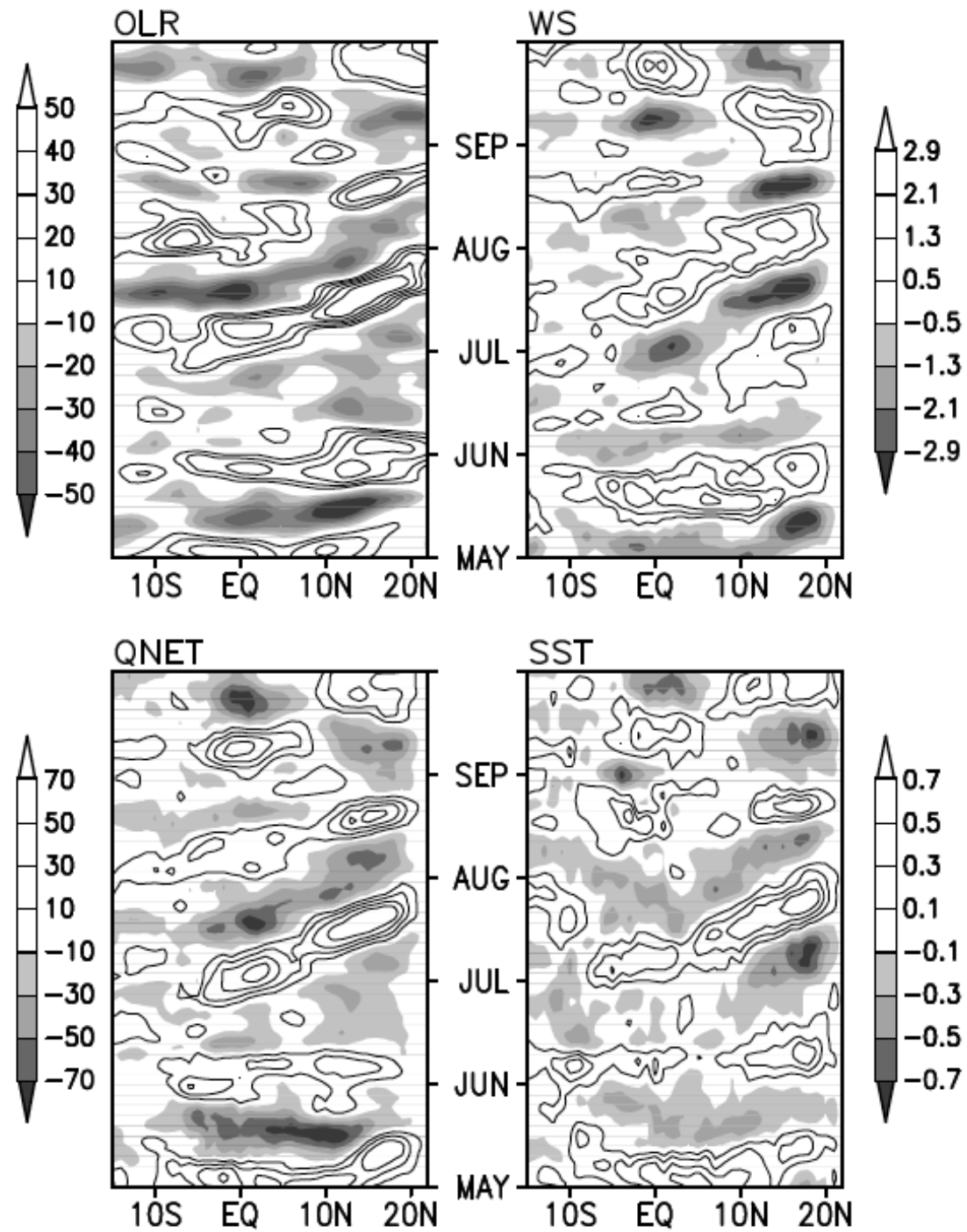
Active phase



Break phase

*Goswami 2003*

*ISO*

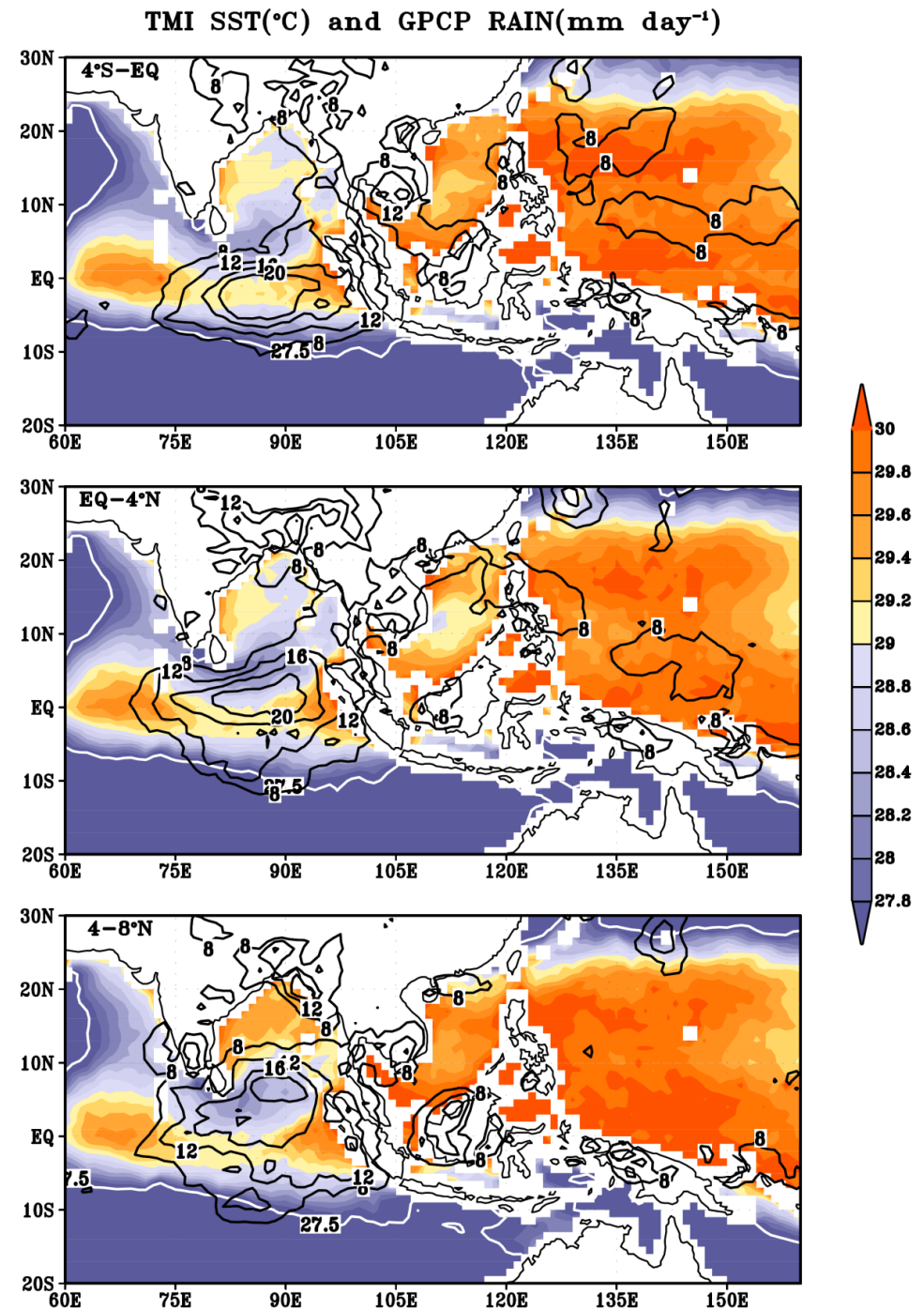


10-80 day  
1998

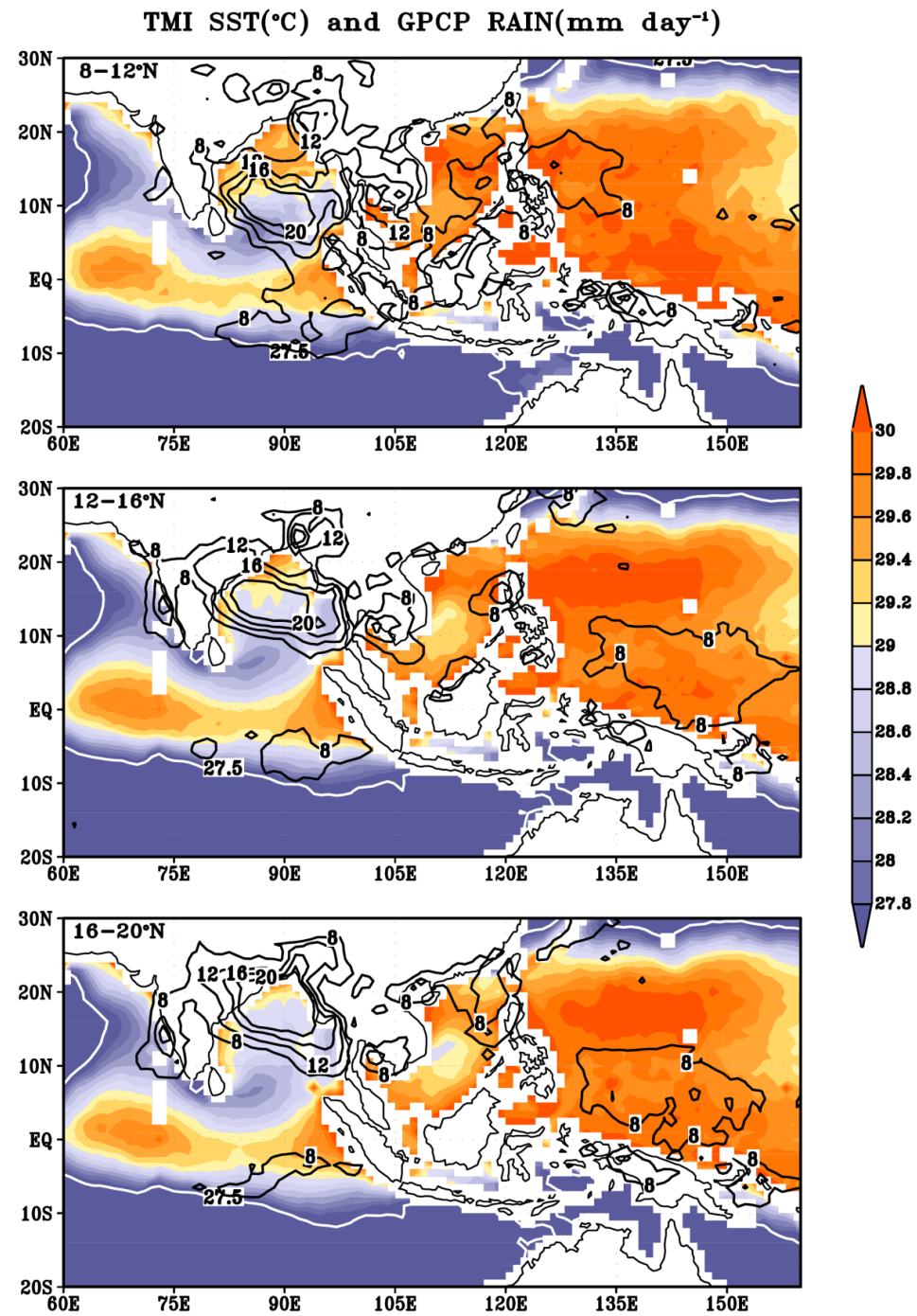


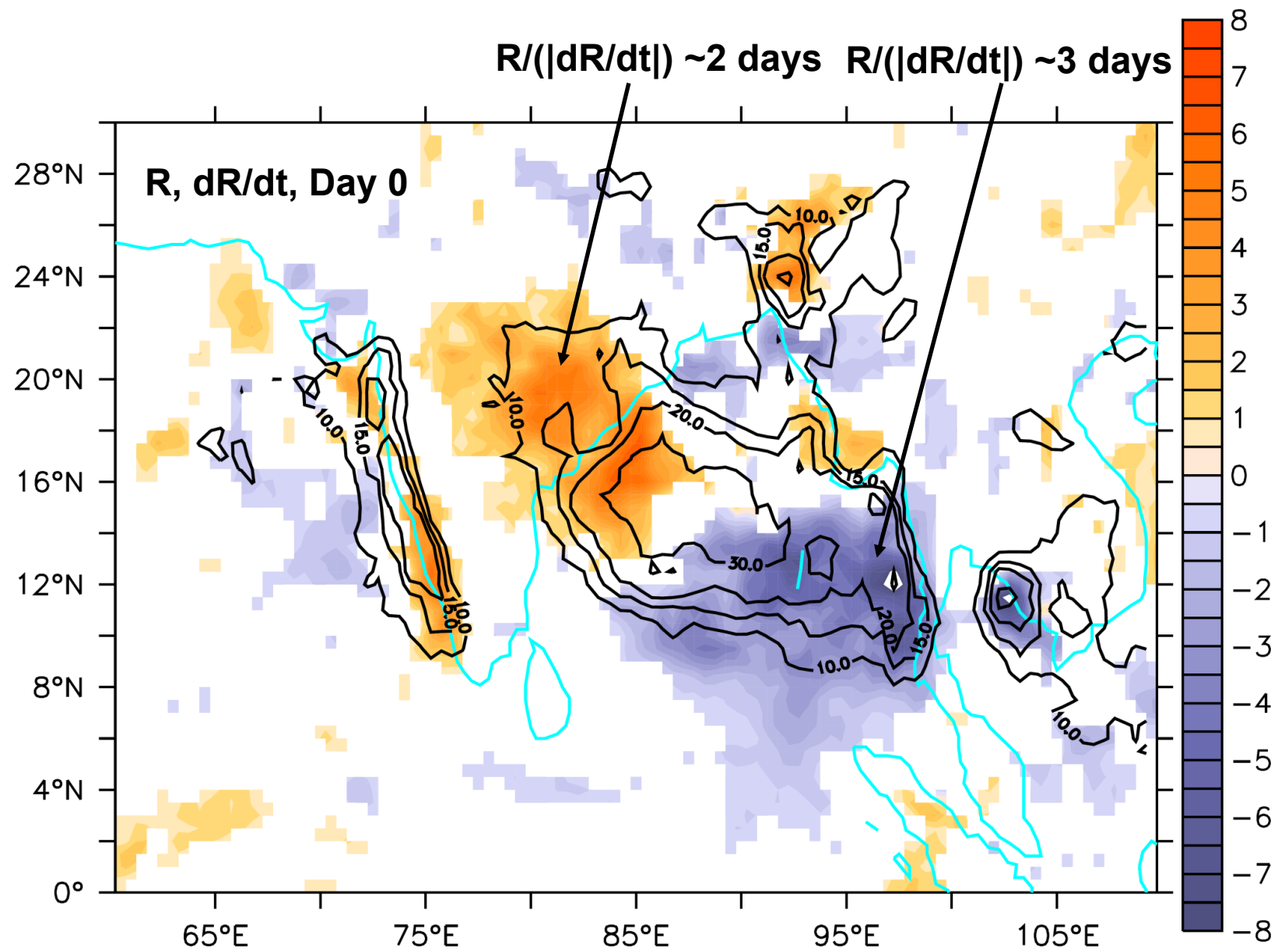
The summer monsoon has 2-3  
“active” and “weak” spells each  
season

Monsoon rainband moves north  
from the equatorial Indian Ocean;  
repeat time ~30 days

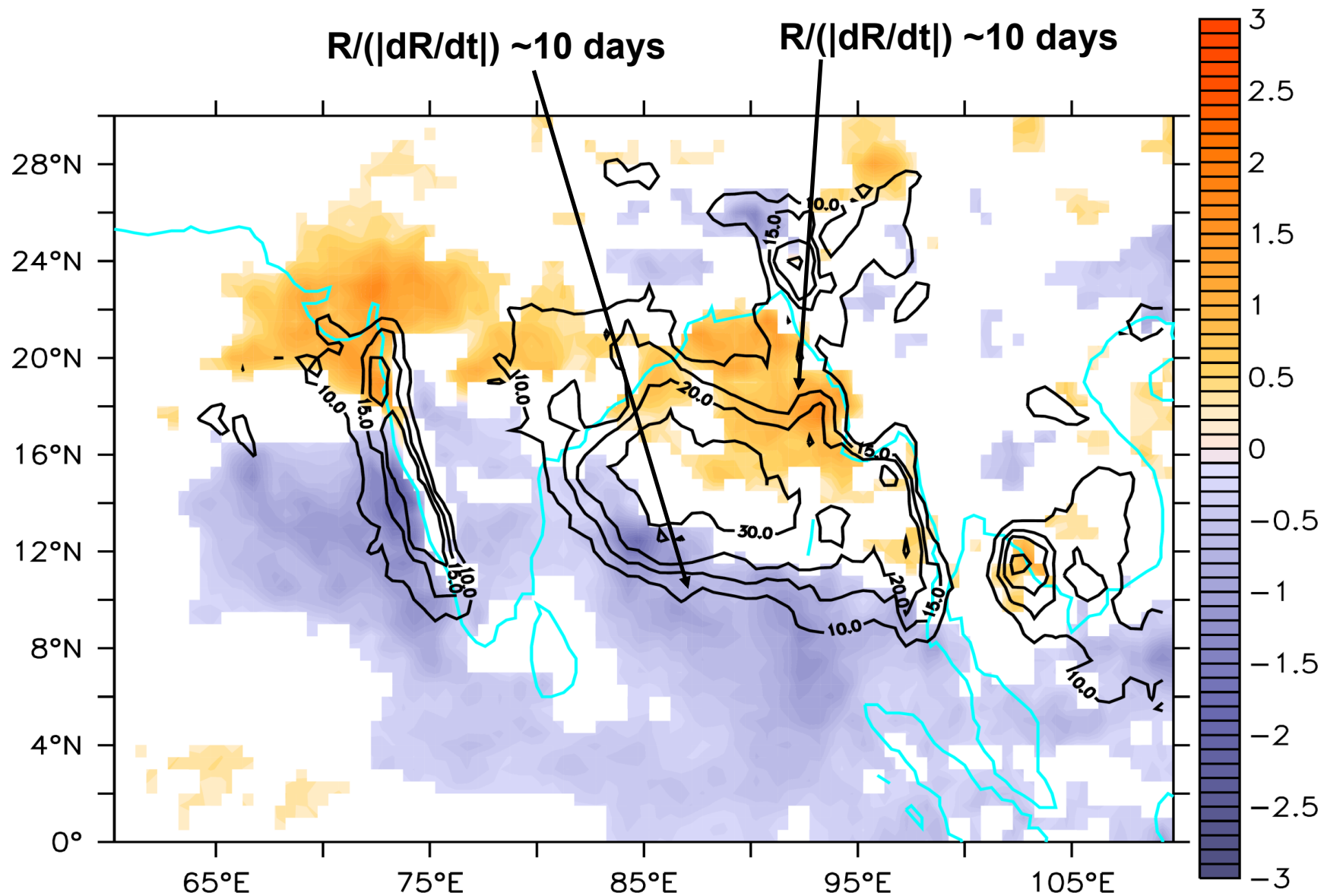


Warm SST to the  
north of the rainband





**Rain grows in northwest and decays in southeast on fast timescale.**

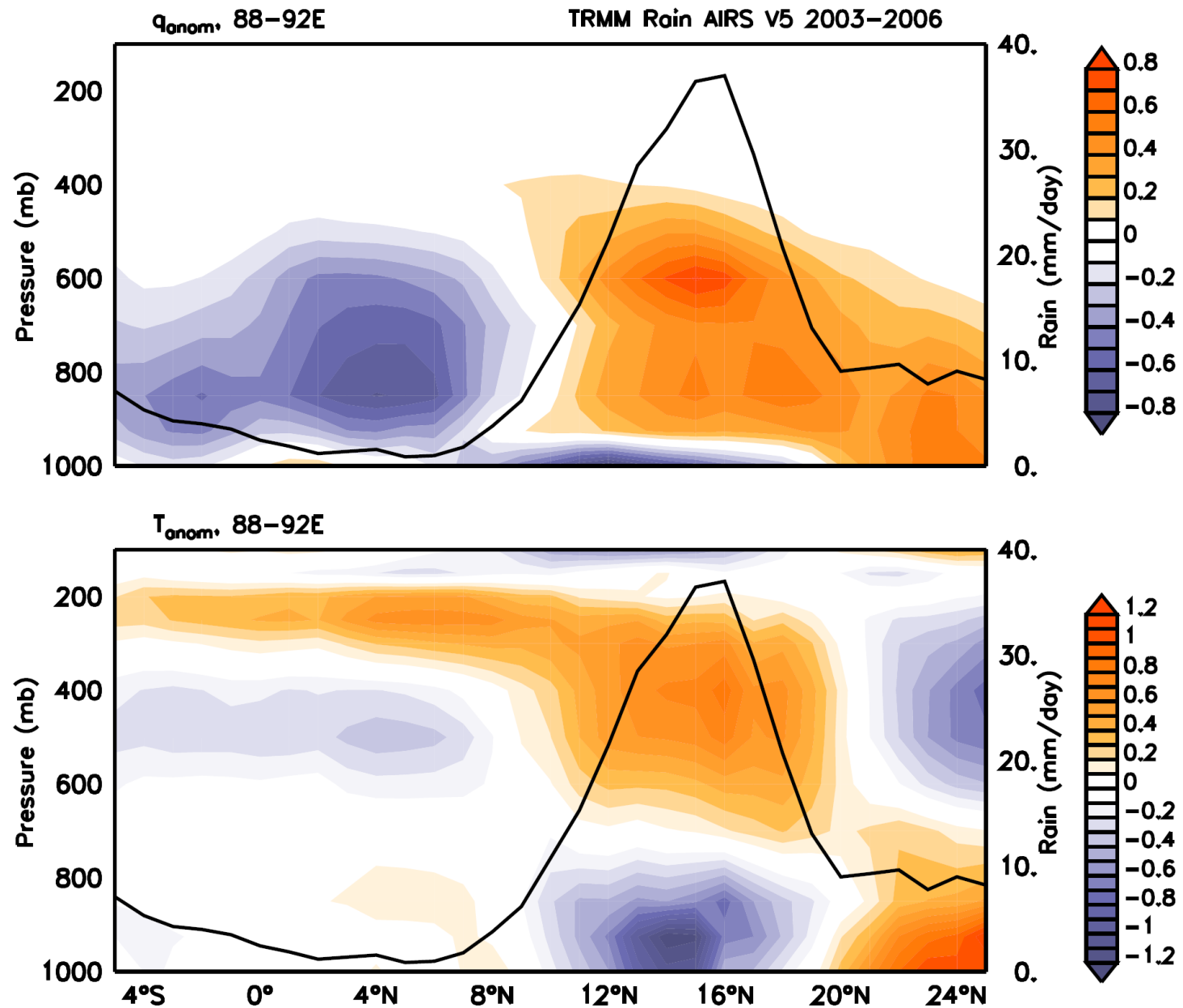


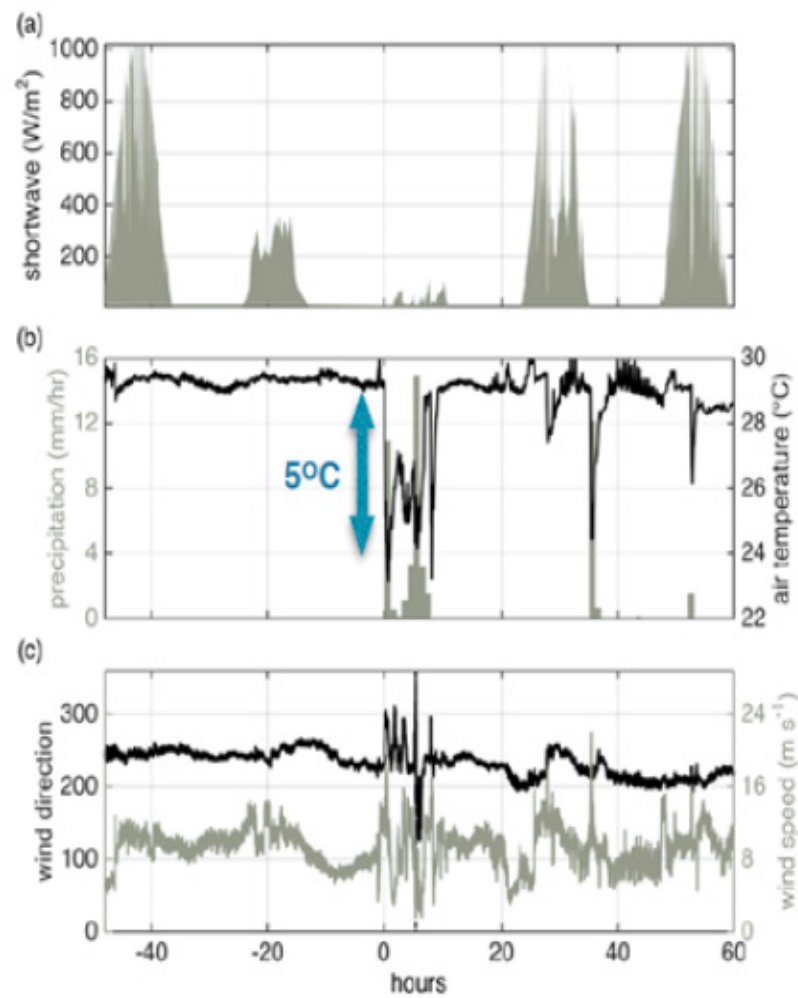
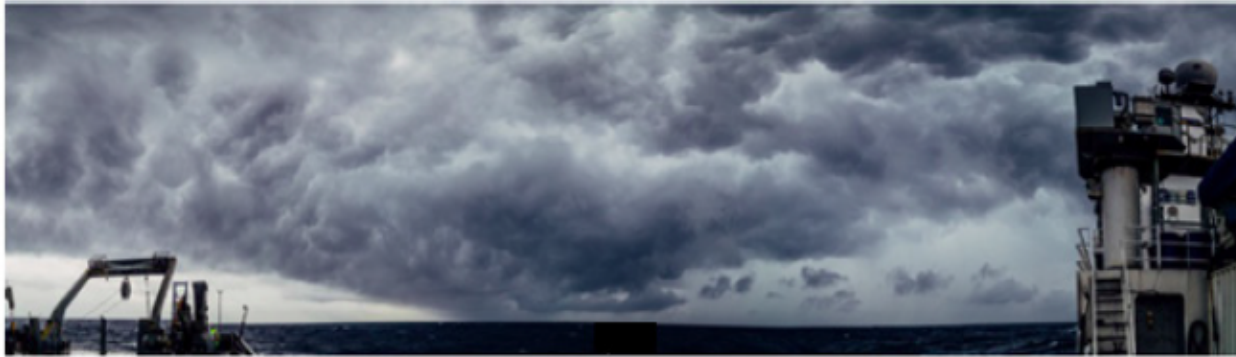
**Movement of rainband parallel to itself on the slow timescale.**

**Monsoon rainfall has two time scales – synoptic and intraseasonal**



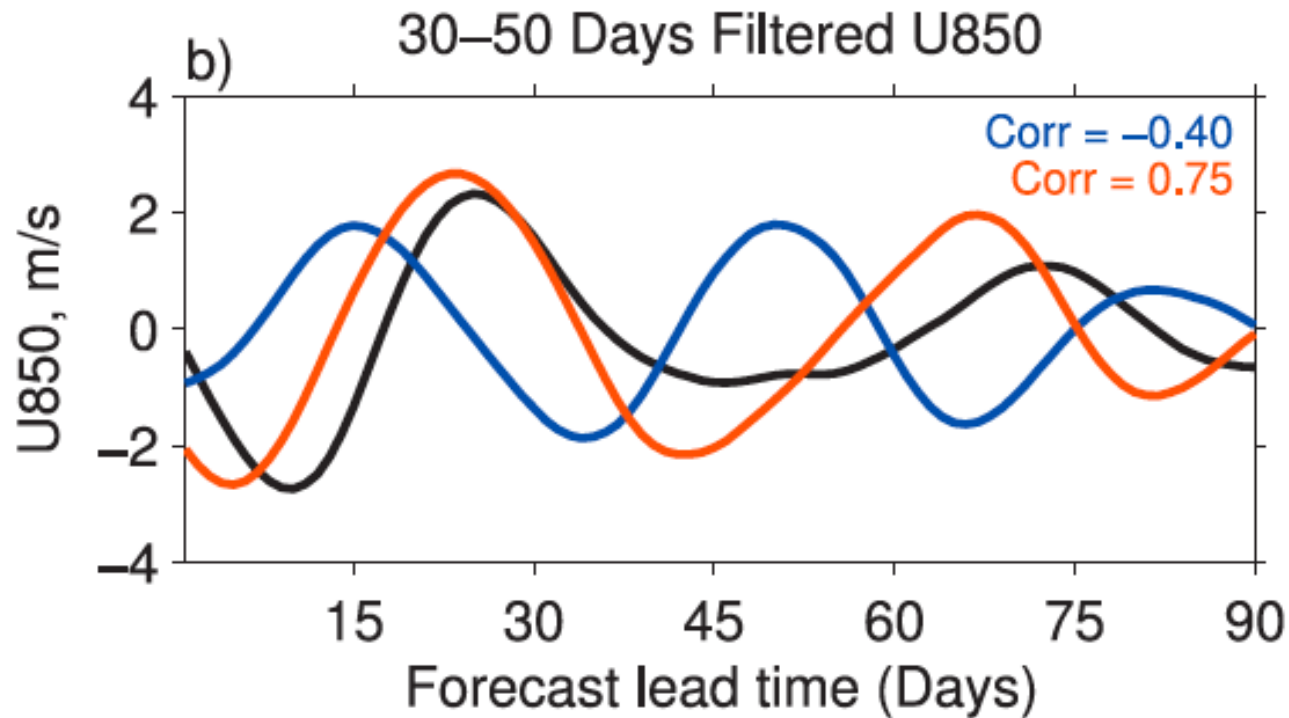
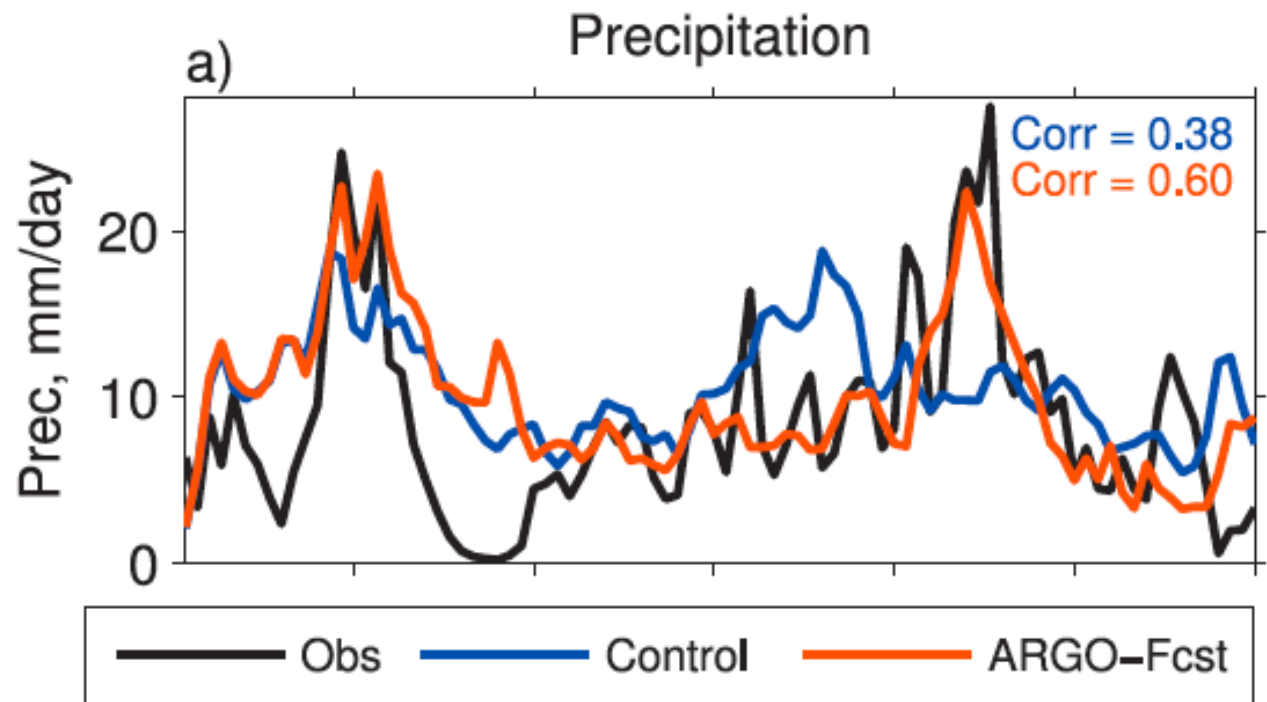
## ISO (2-90 day) $q$ , $T_a$





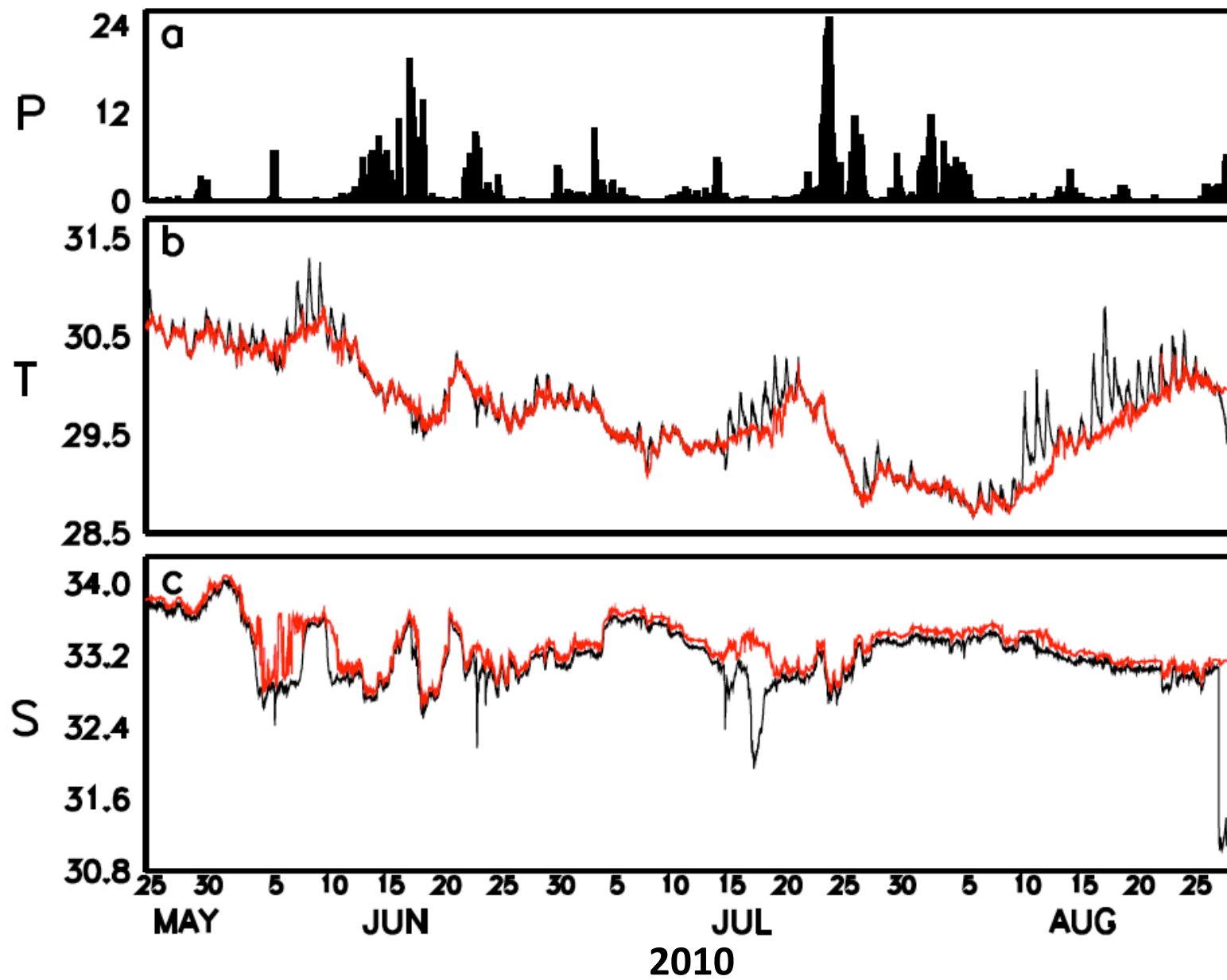
**26 August 2015** *Roger Revelle*

*Courtesy: Emily and San*



*TN Krishnamurti 2007*  
*DeMott 2011*

18N 89.5E

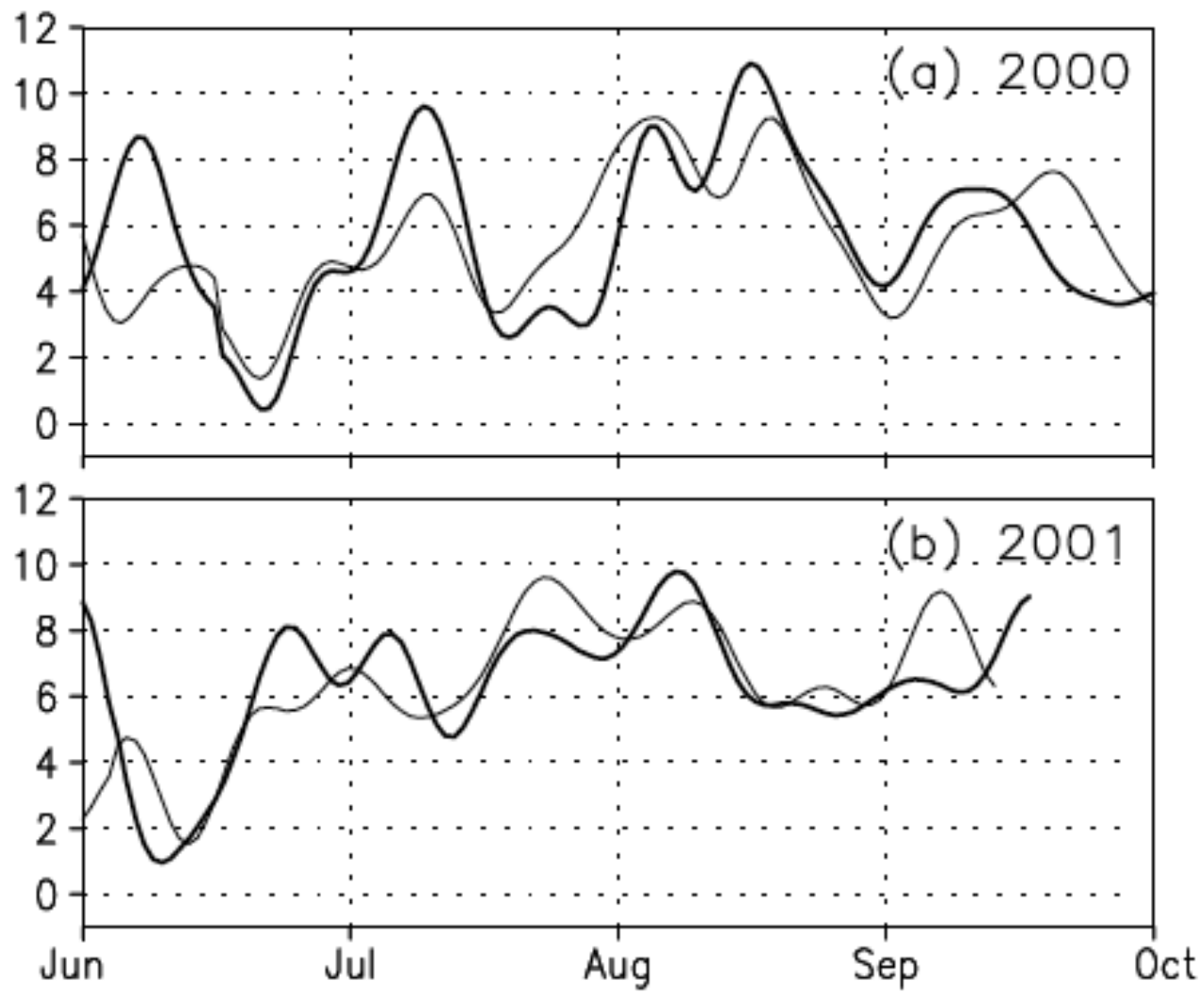






**7<sup>th</sup> century cave temple Trichy**

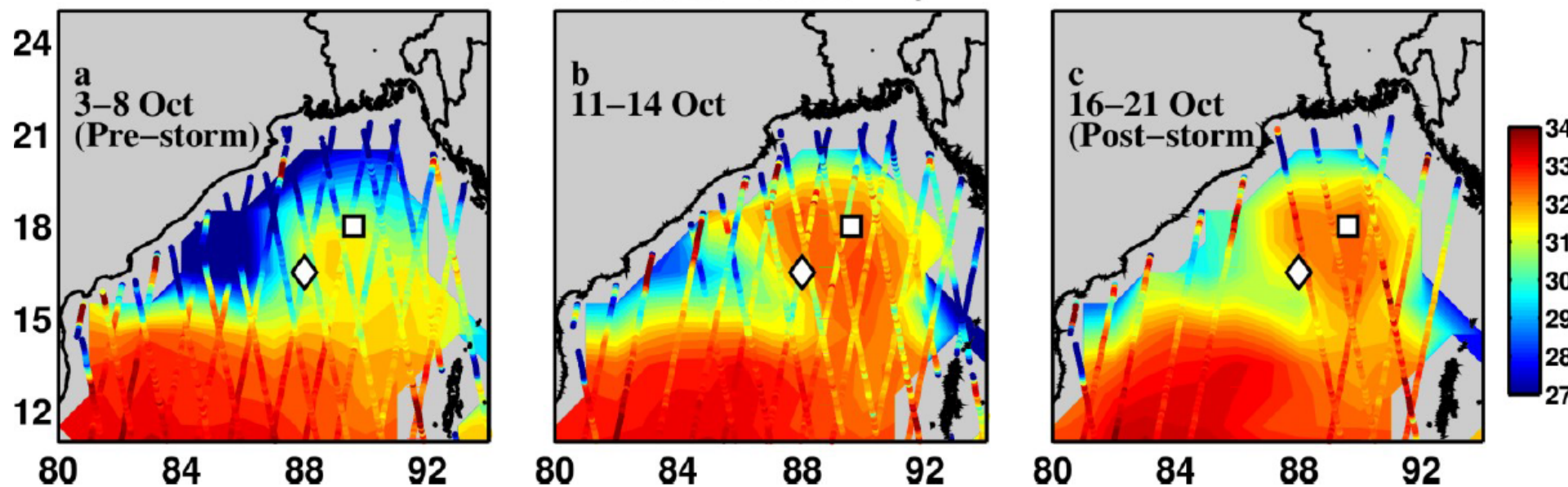




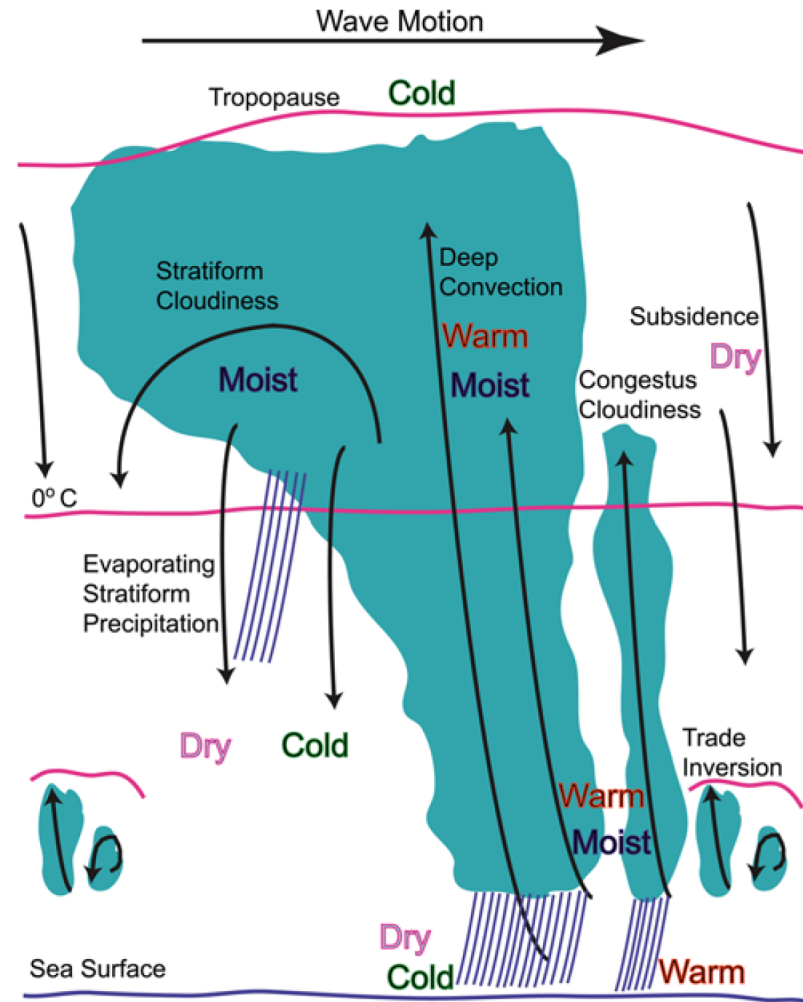
*Goswami 2003*



### Sea Surface Salinity







Kiladis et al. 2008,  
Rev. Geophys.

**Figure 19.** The hierarchy of cloudiness, temperature, and humidity within CCEWs, valid from MCS to MJO scales. Wave movement is from left to right (adapted from *Johnson et al.* [1999], *Straub and Kiladis* [2003c], and *Khouider and Majda* [2008]).

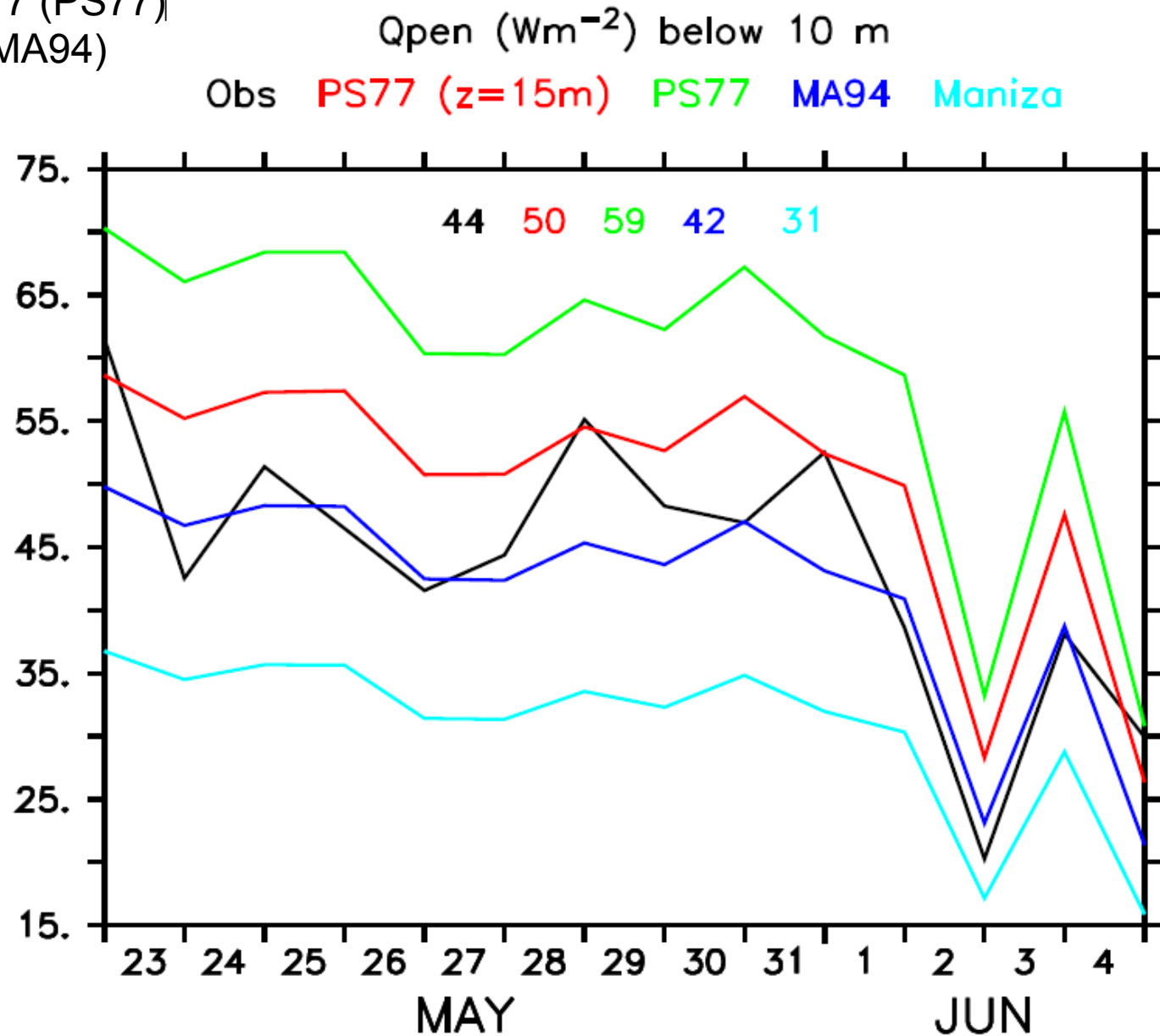
## Q<sub>open</sub>

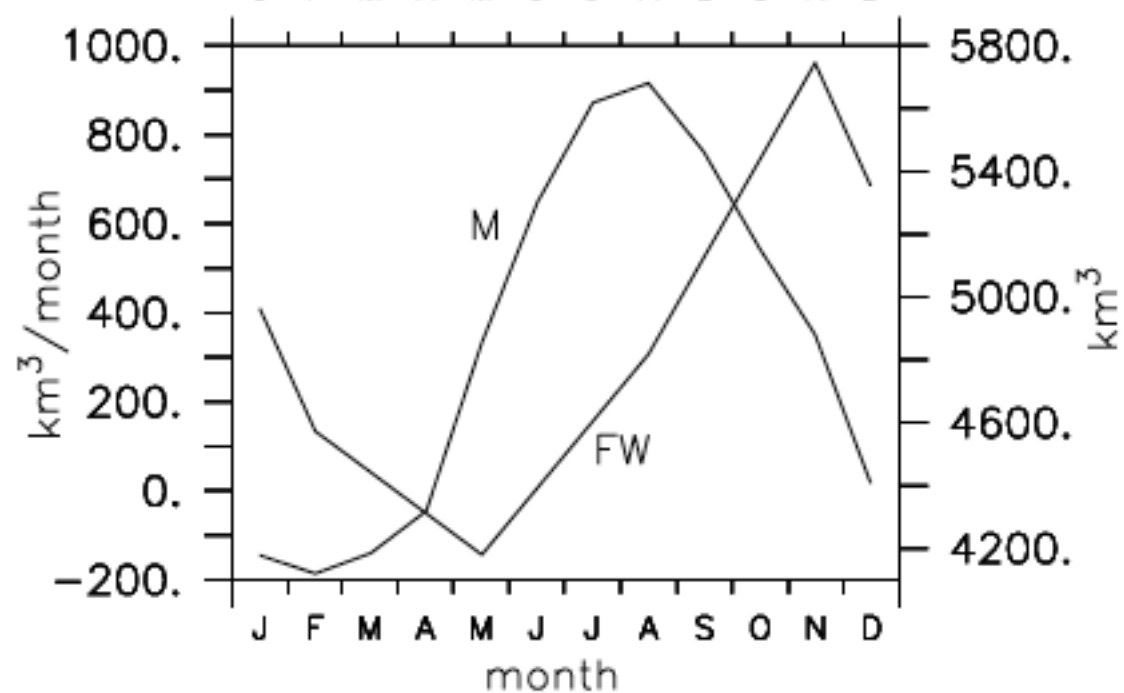
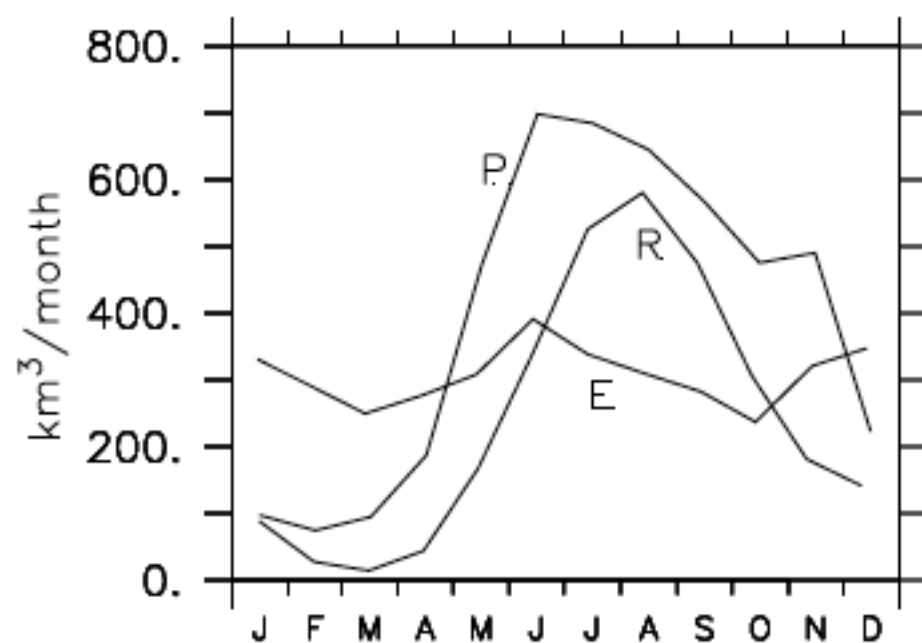
Paulson-Simpson 1977 (PS77)

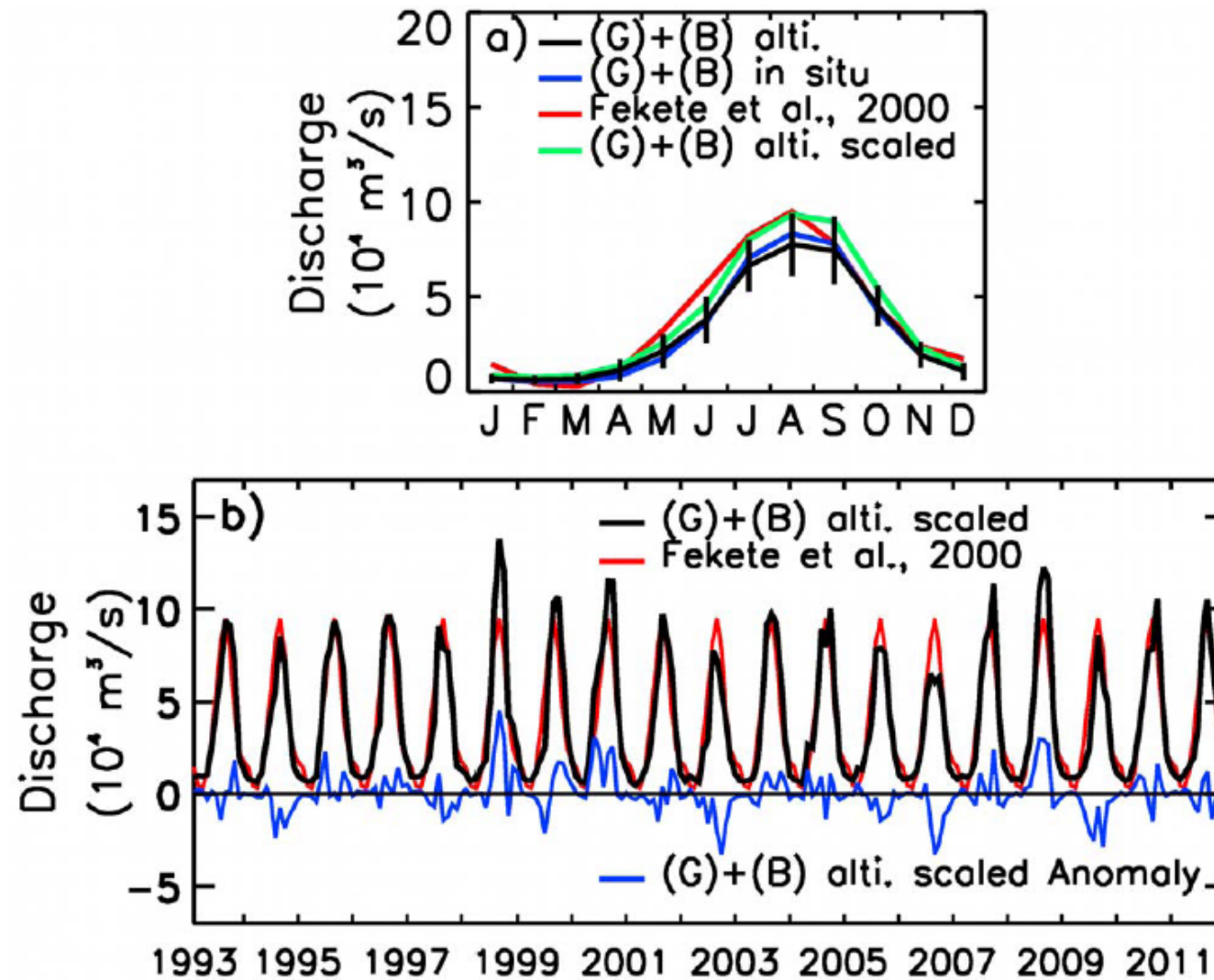
Morel-Antoine 1994 (MA94)

Manizza et al., 2005

Weekly SeaWIFS chl.

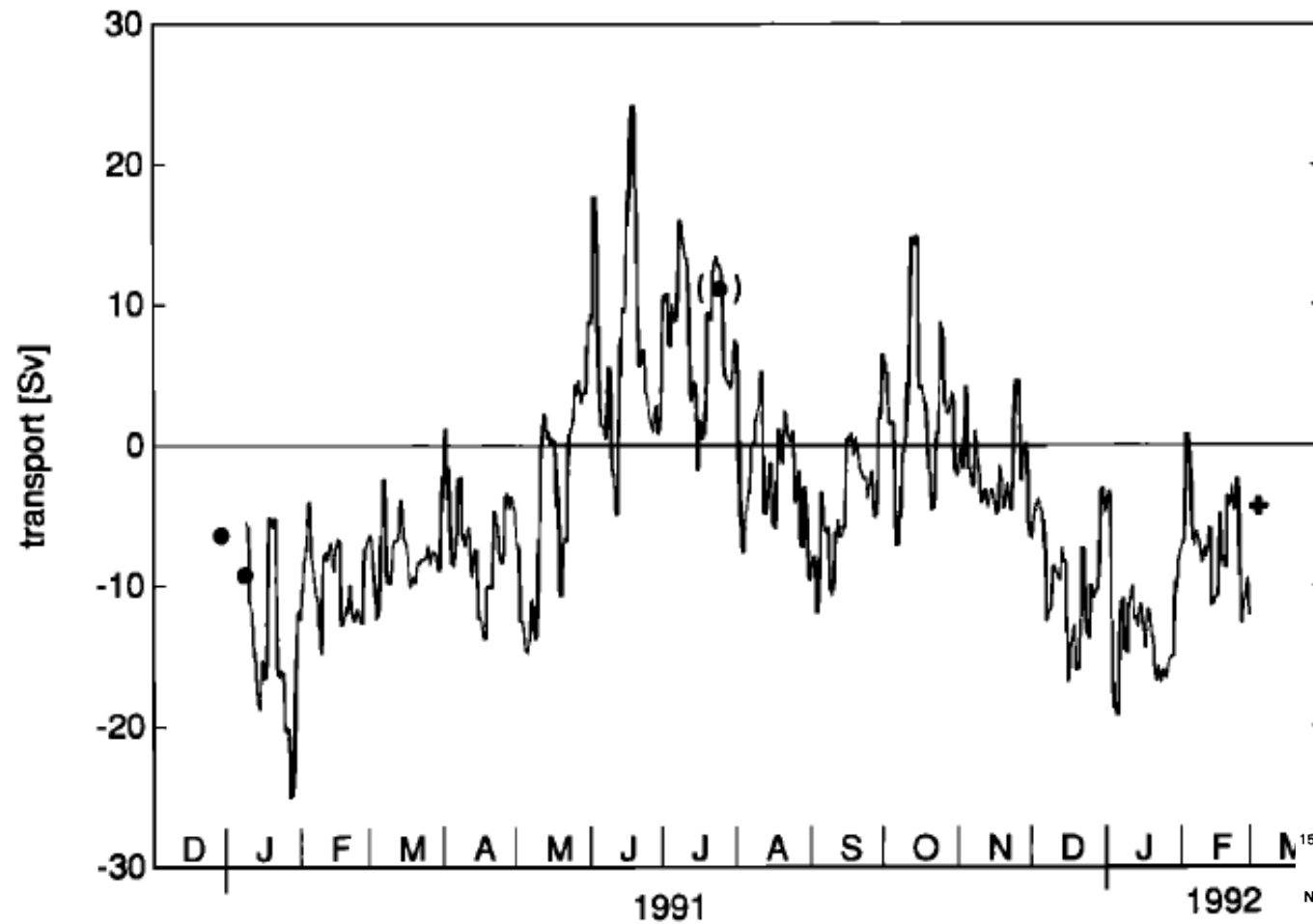




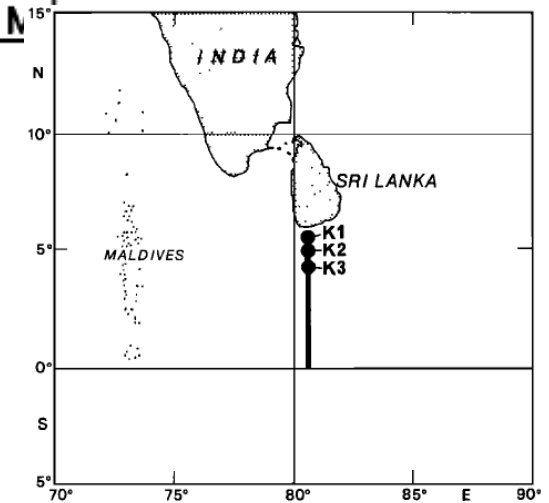


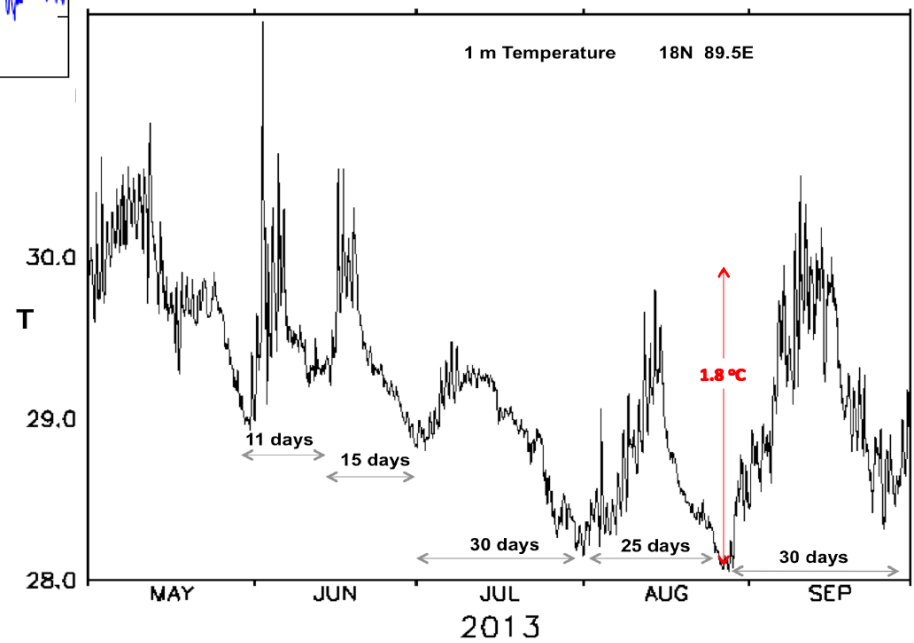
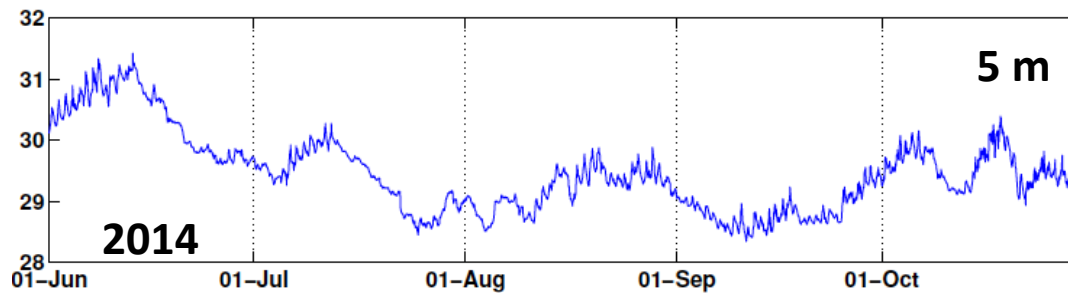
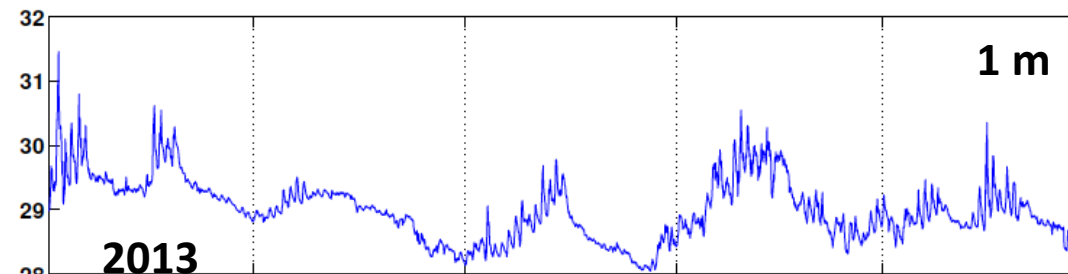
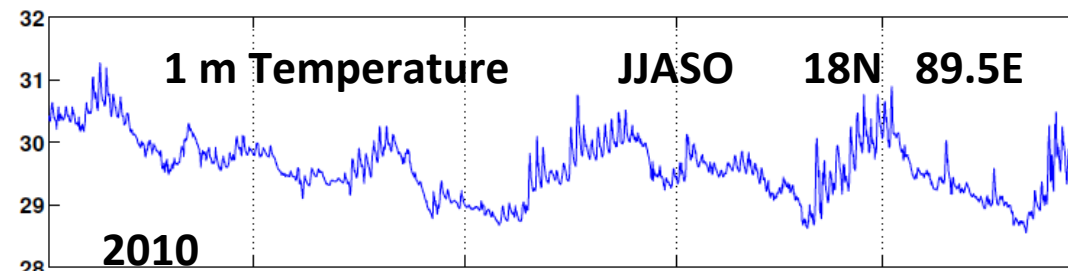
G-B-M discharge



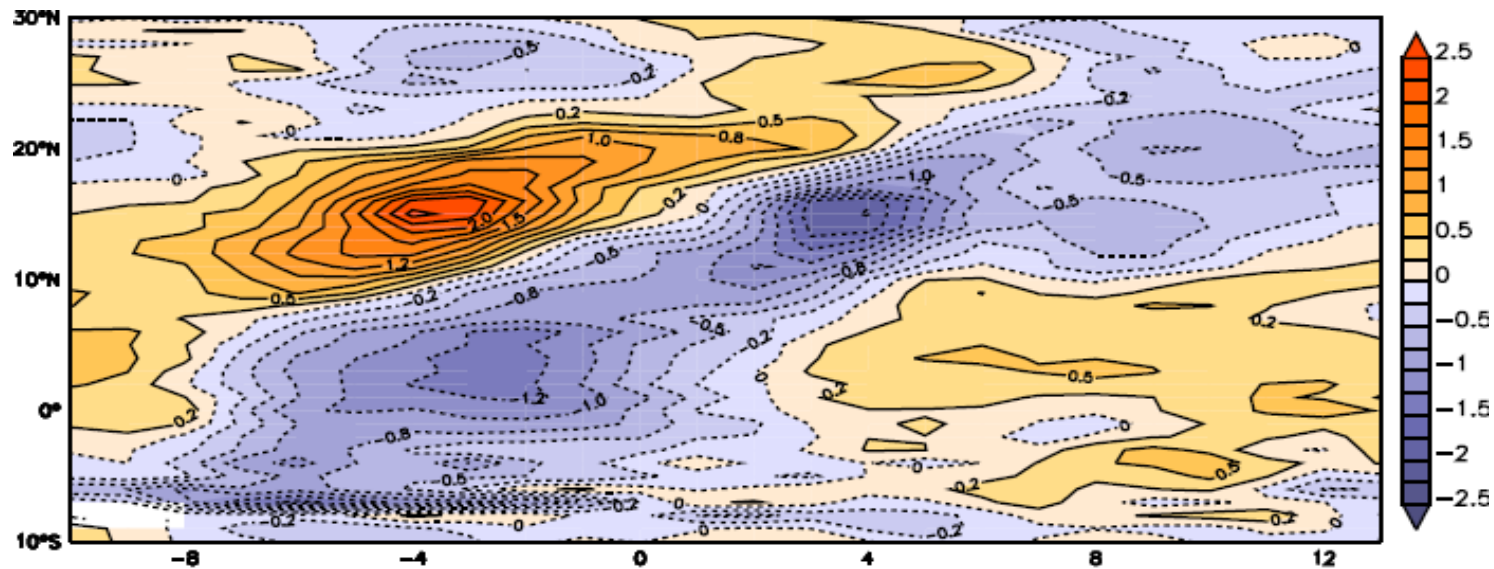


Schott *et.al* 1994

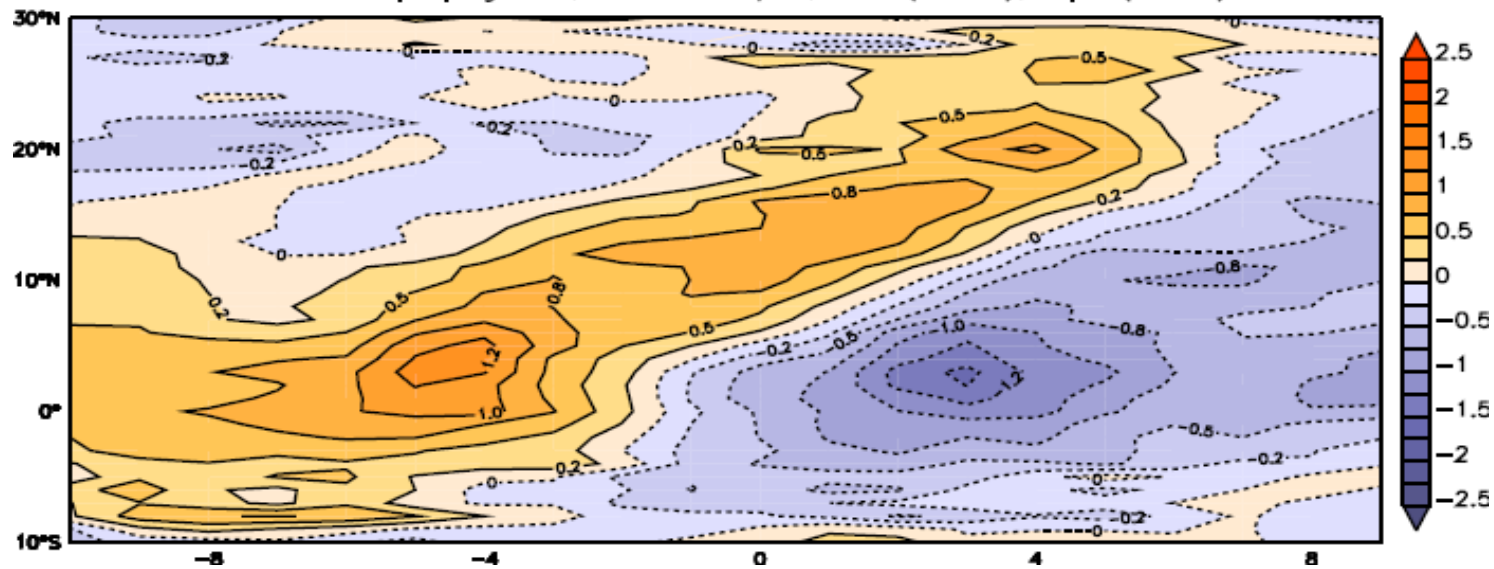




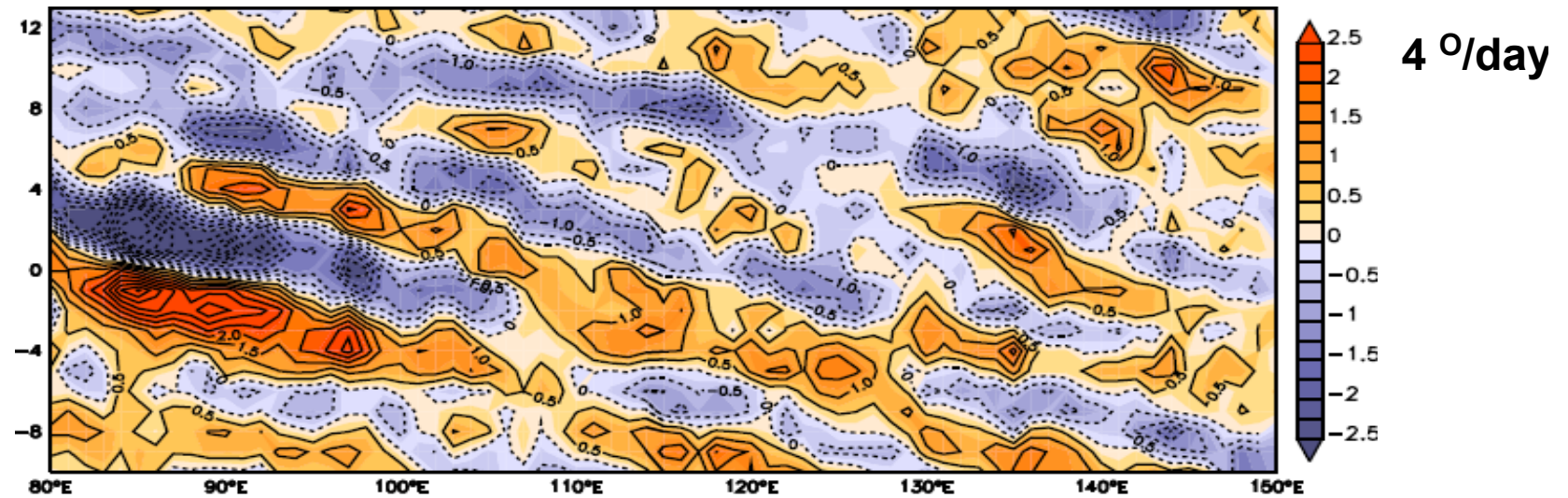
## 7-day R, dR/dt CBay, EqWestPac



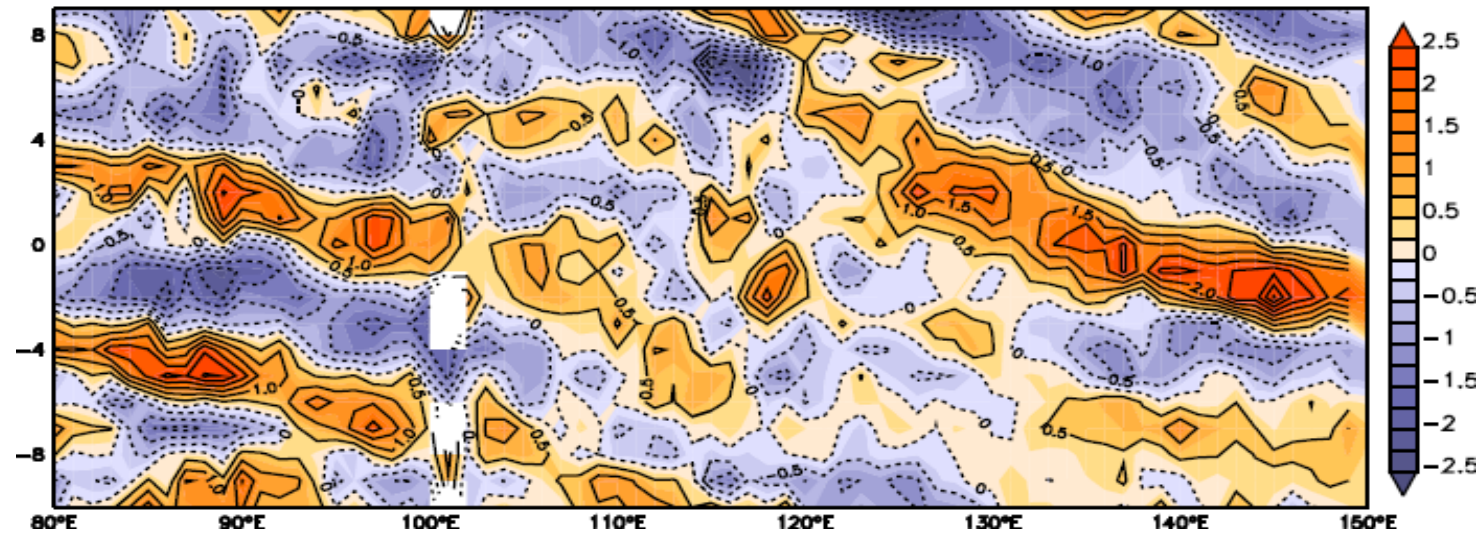
Northward propagation, smooth dR/dt, CBay(above), EqWP(Below)



## $dR/dt$ Cbay, EqWestPac

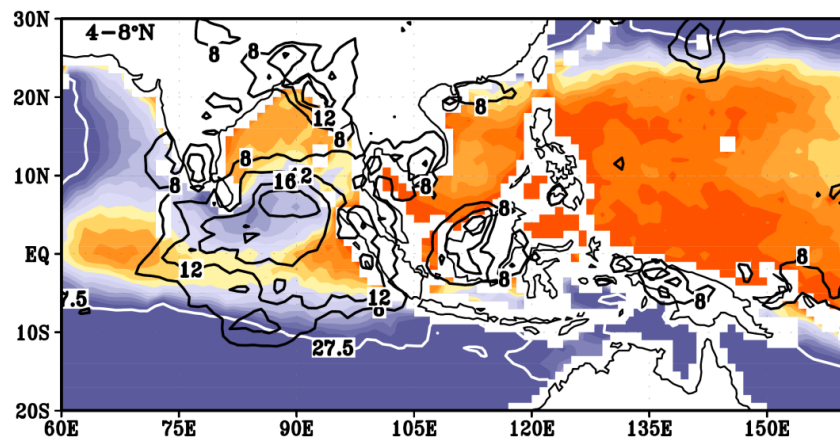
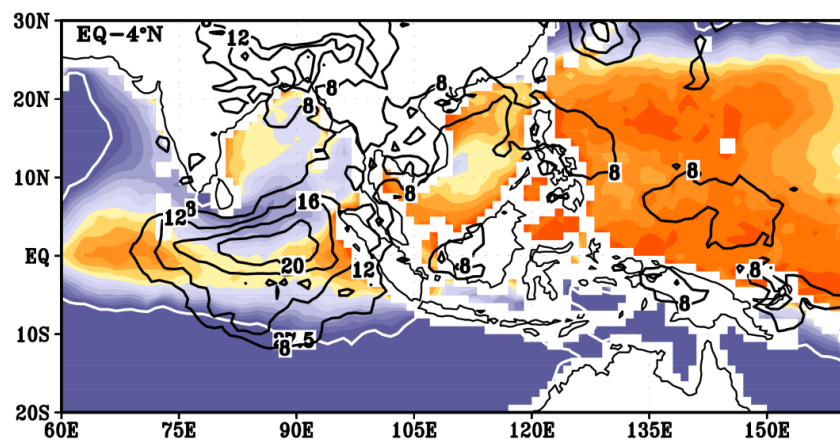
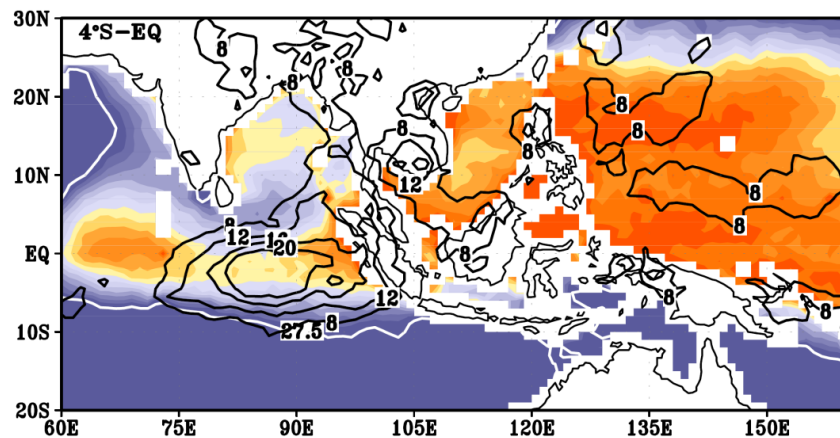


Westward propagation, nonsmooth  $dR/dt$ , CBAY(above), EqWP(Below)

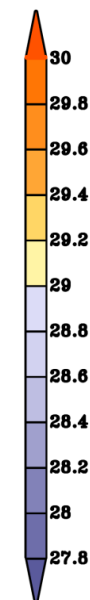
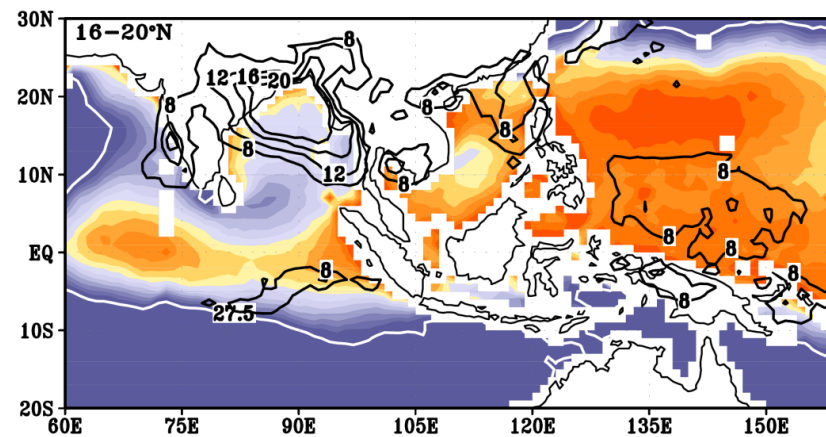
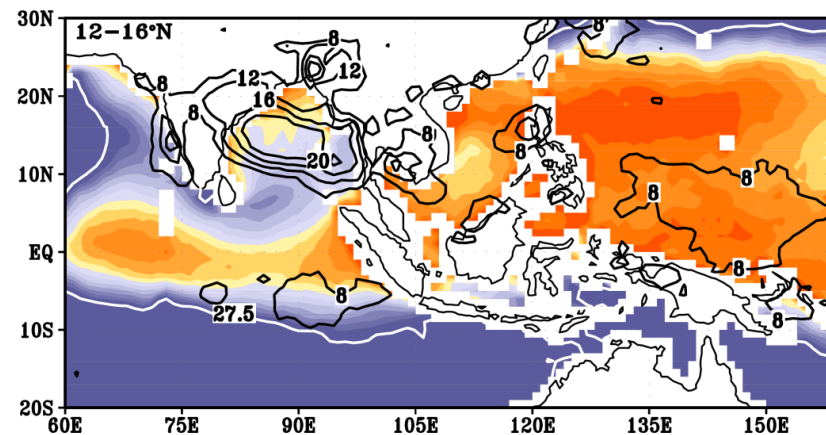
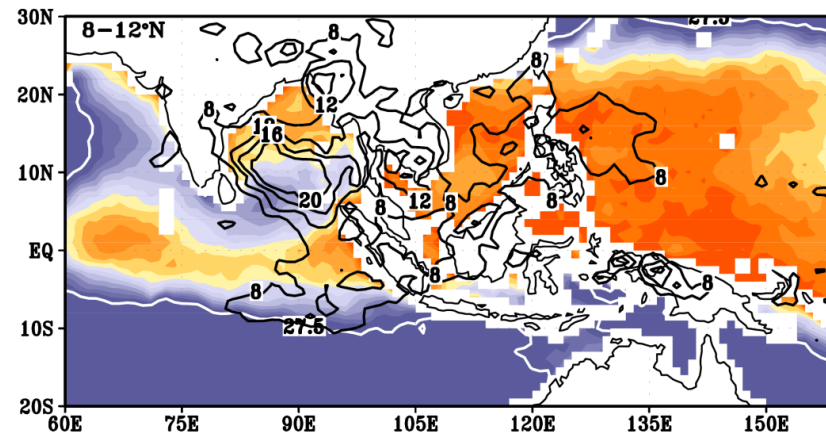




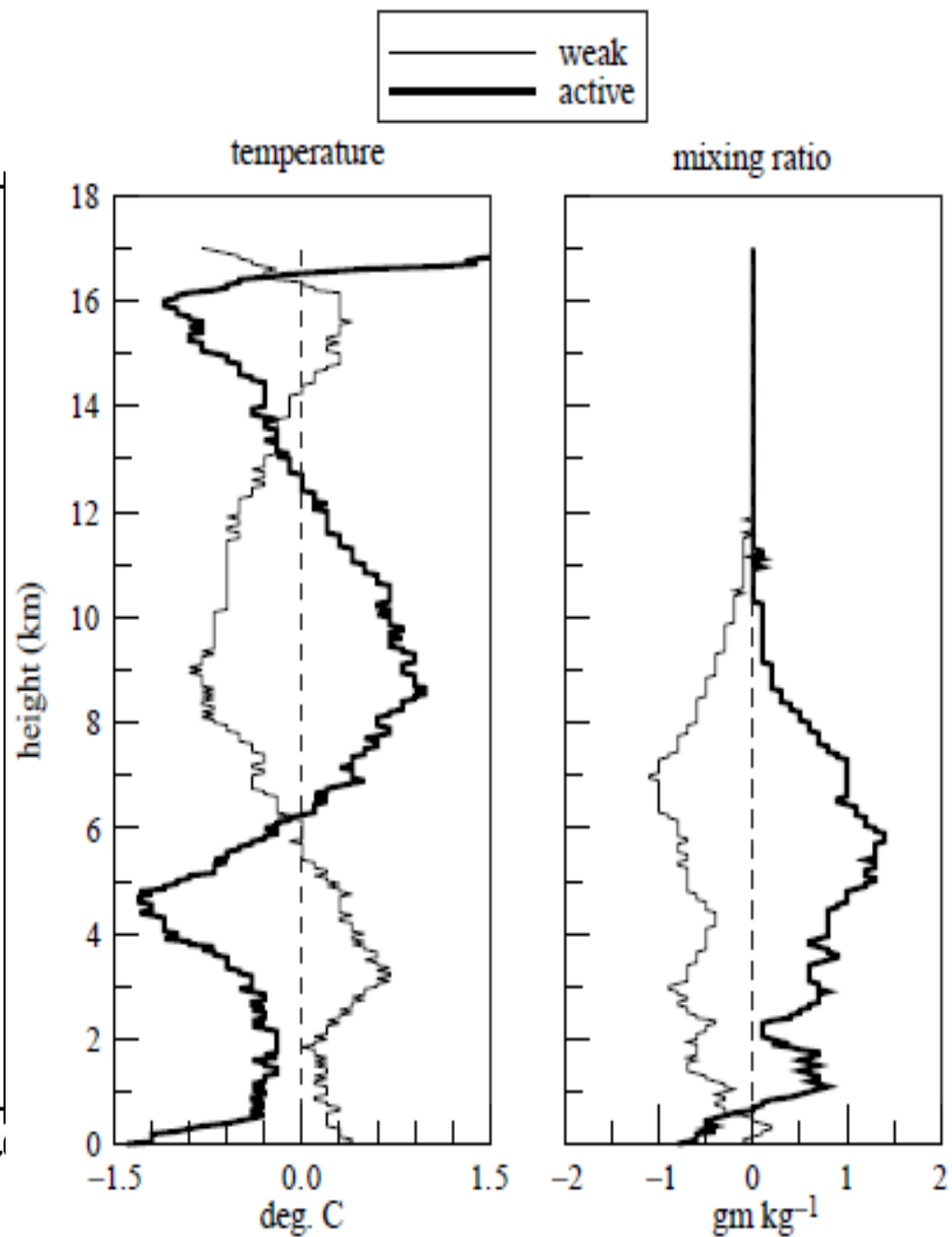
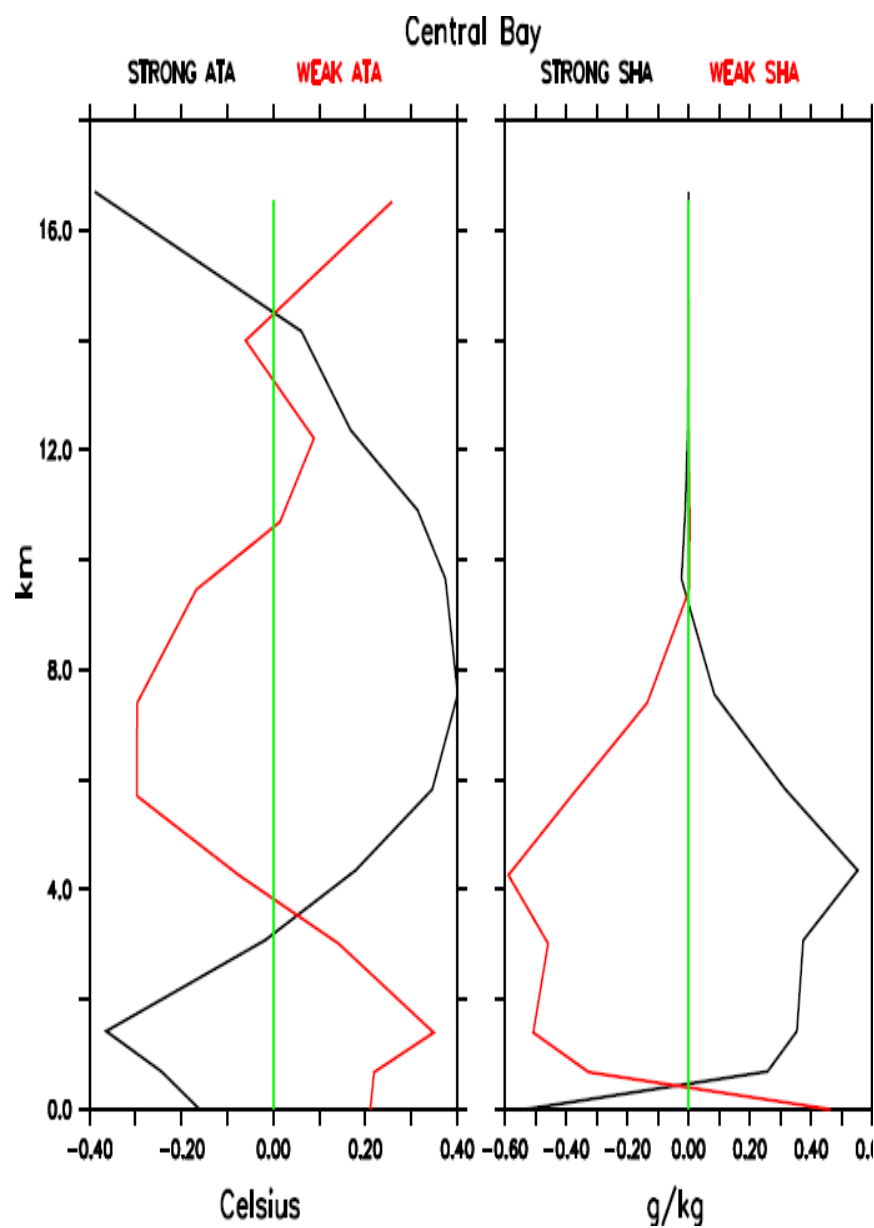
TMI SST( $^{\circ}\text{C}$ ) and GPCP RAIN( $\text{mm day}^{-1}$ )

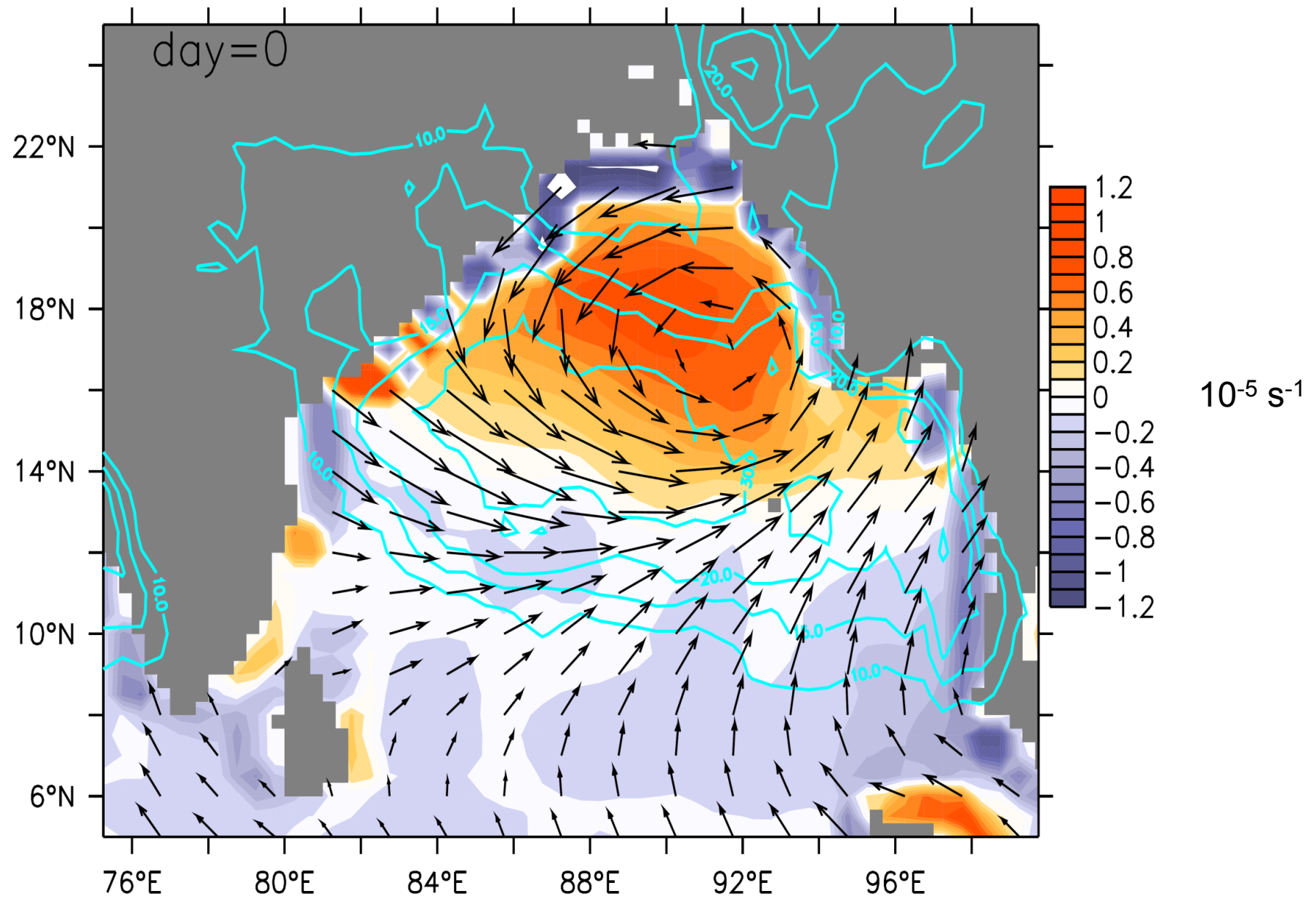


TMI SST( $^{\circ}\text{C}$ ) and GPCP RAIN( $\text{mm day}^{-1}$ )



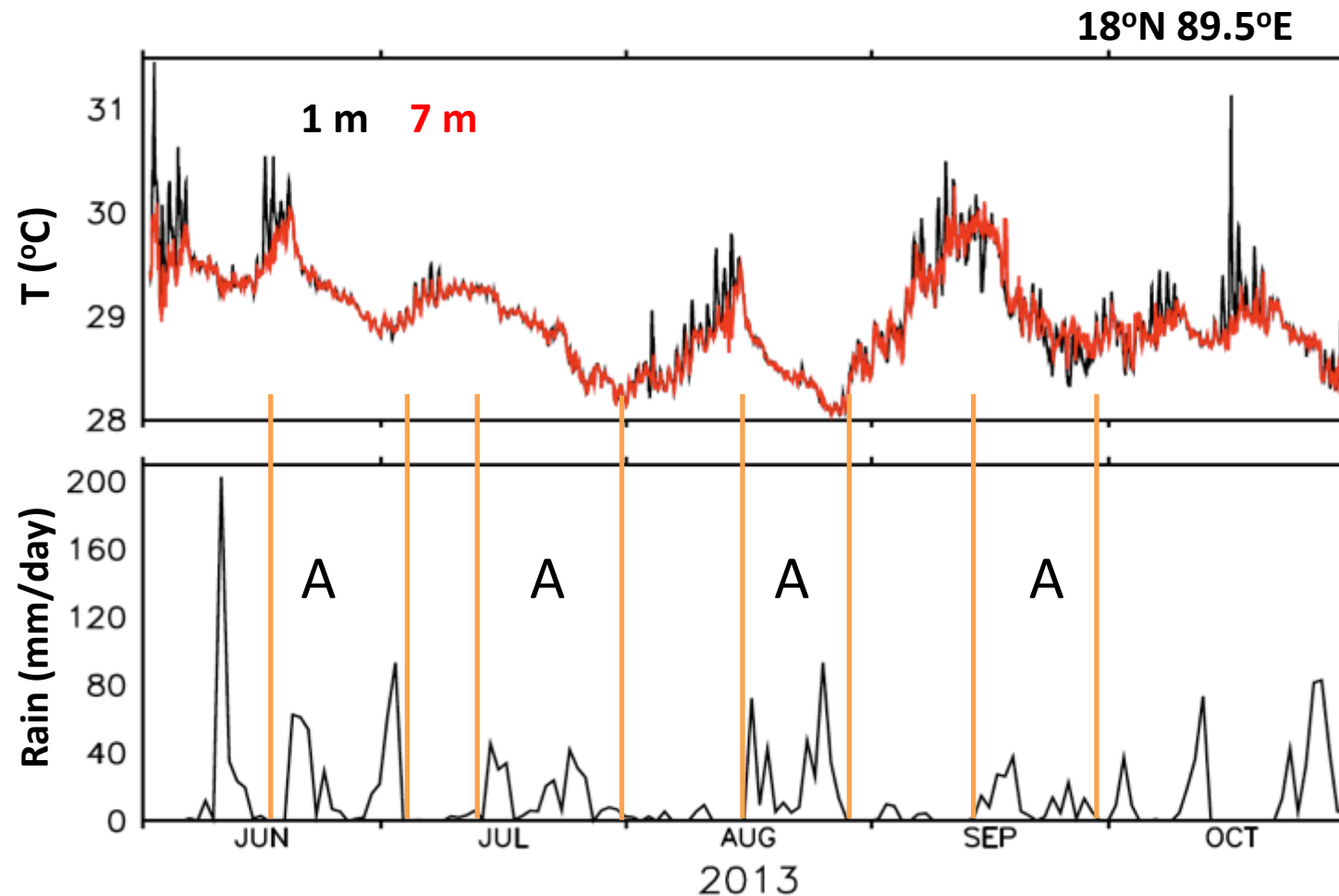
# AIRS





**Maximum vorticity anomaly is 3-4 degrees north of the maximum rain**

## SST and monsoon rainfall



**SST cools during active (cloudy, windy) spells**

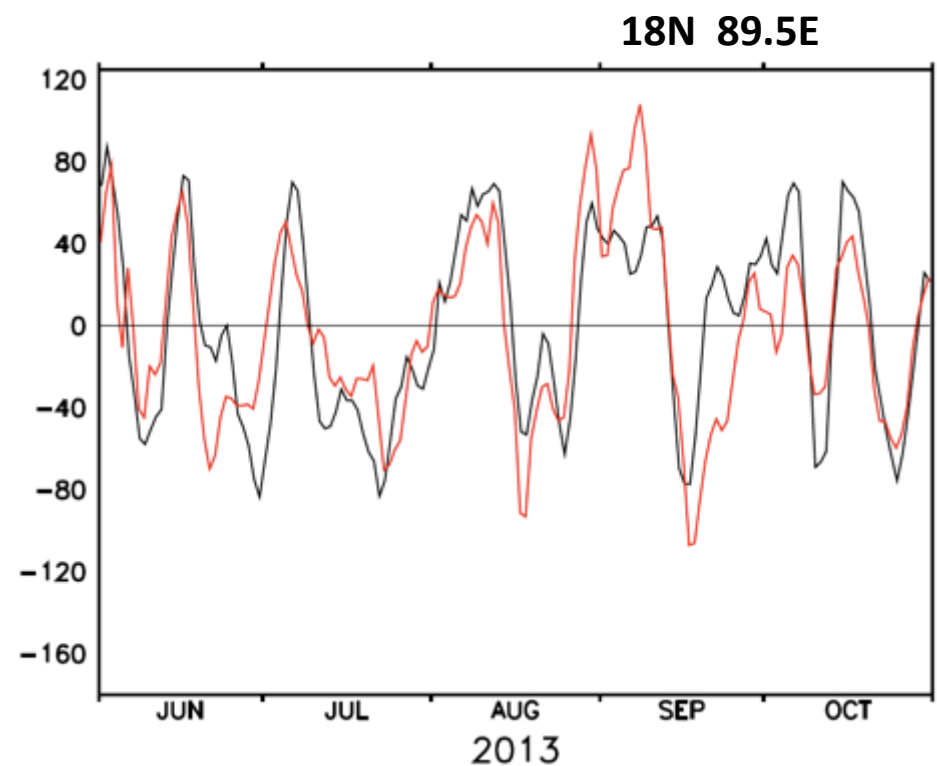
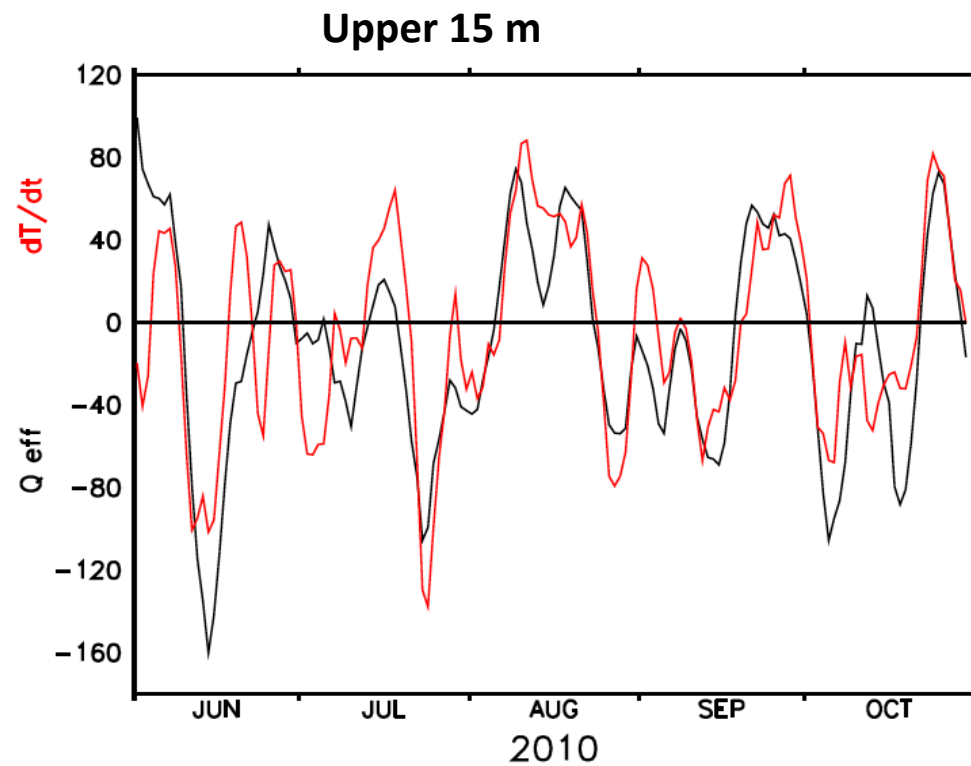
**SST oscillations influence active-break spells of the summer monsoon**

*Sengupta Ravichandran 2001, Vecchi Harrison 2002, De Mott 2011*

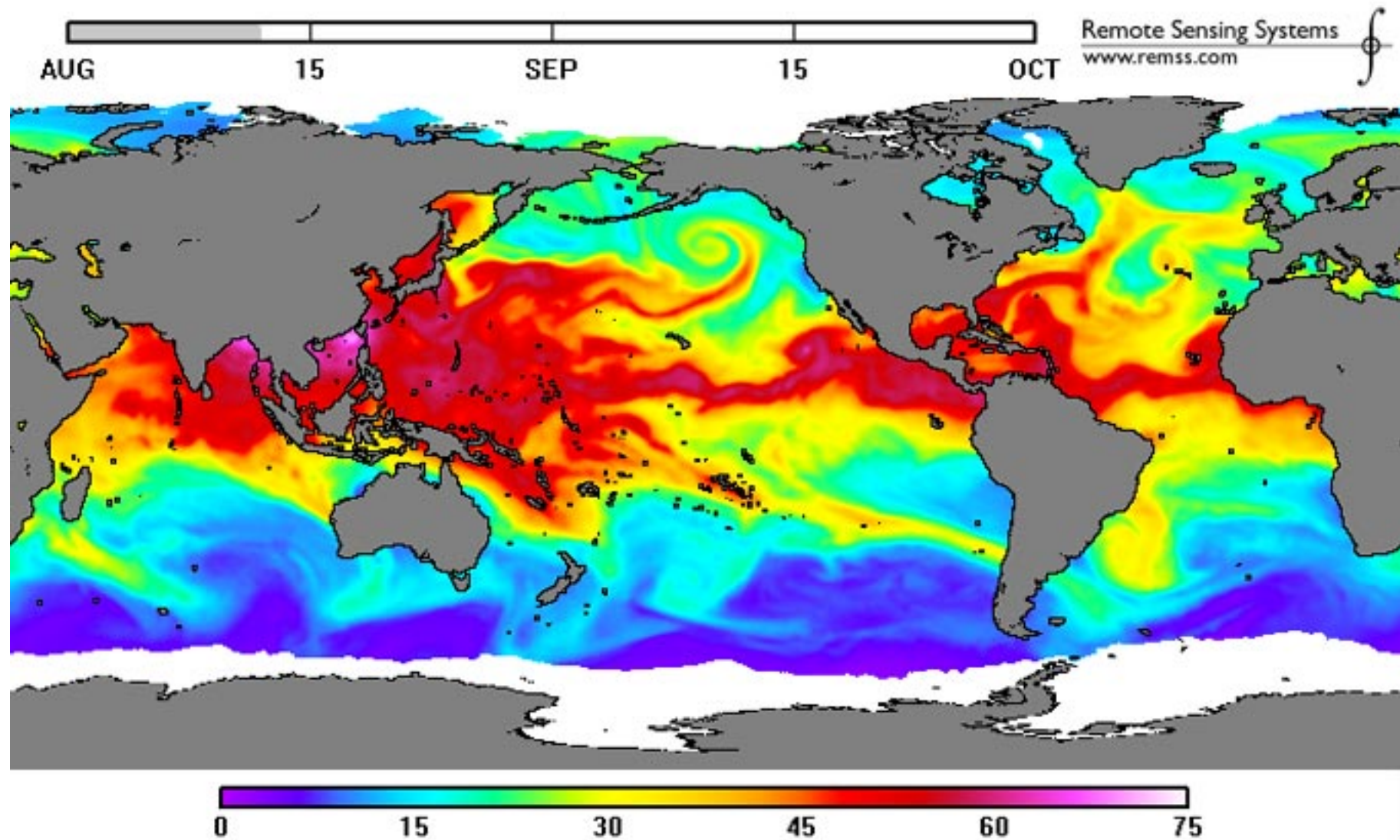


## Heat balance equation

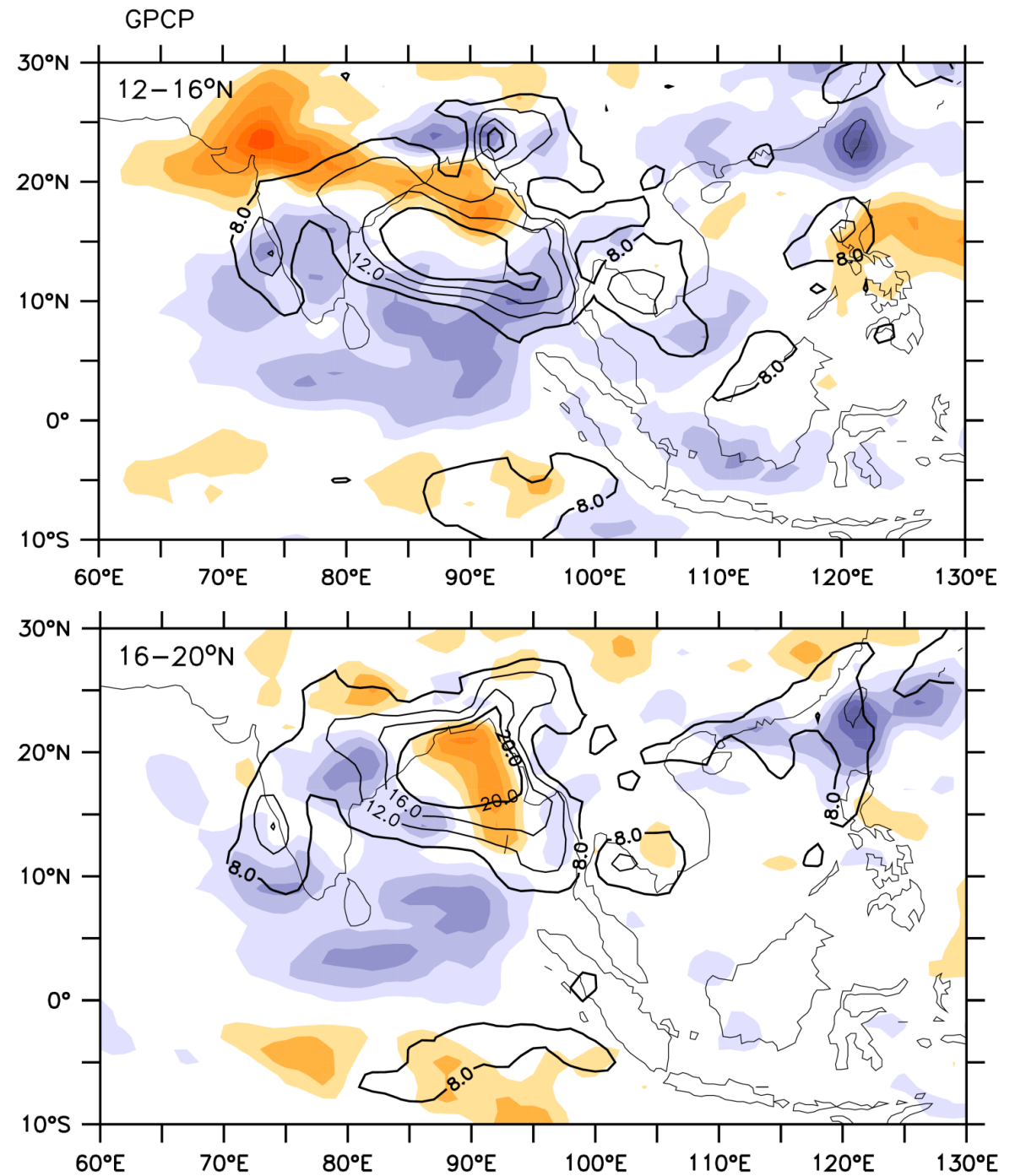
$$\frac{\partial SST}{\partial t} = \frac{1}{\rho C_p H} Q - u \frac{\partial SST}{\partial x} - v \frac{\partial SST}{\partial y} - w \frac{\partial SST}{\partial z} + \text{Mixing}$$



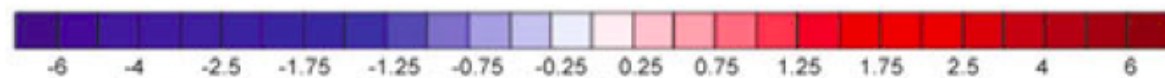
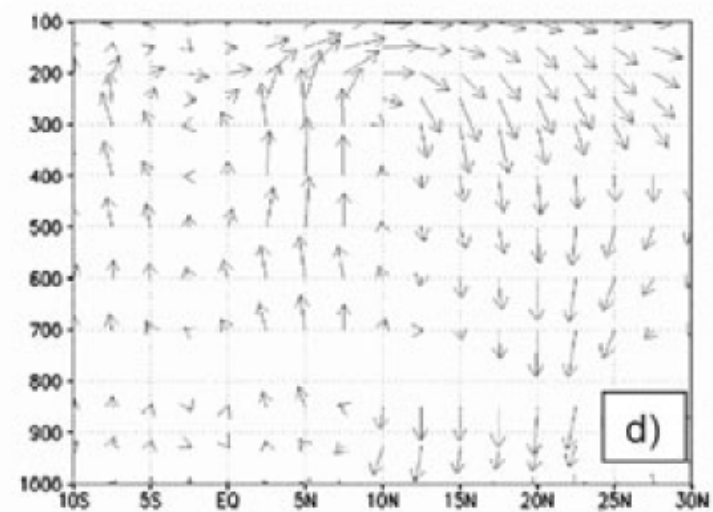
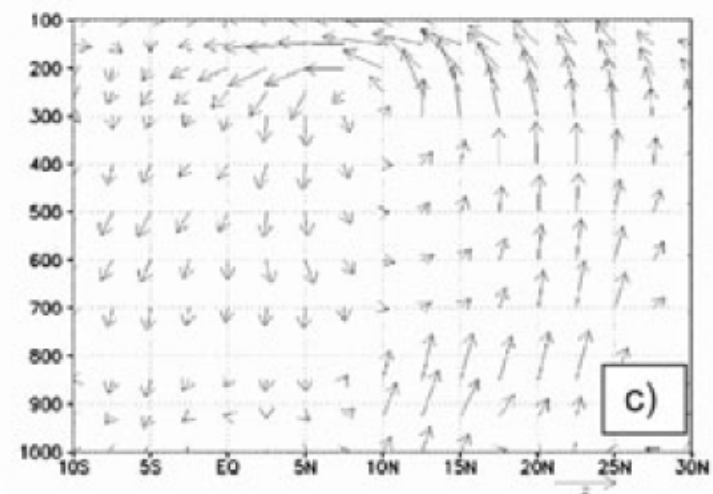
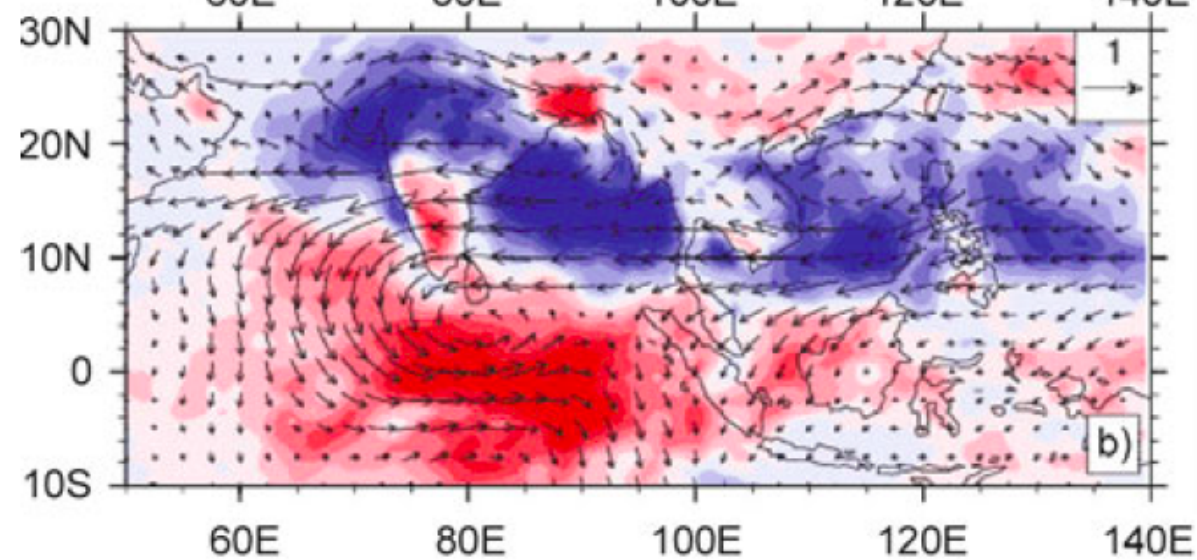
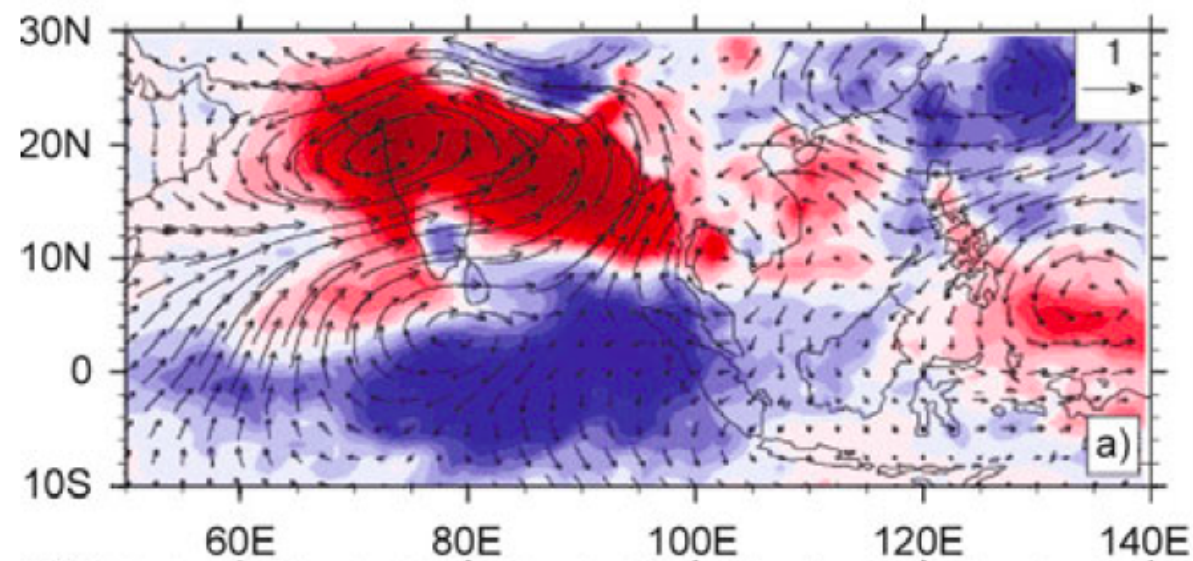
## Column water vapour (mm) from SSM/I 12 August 2005



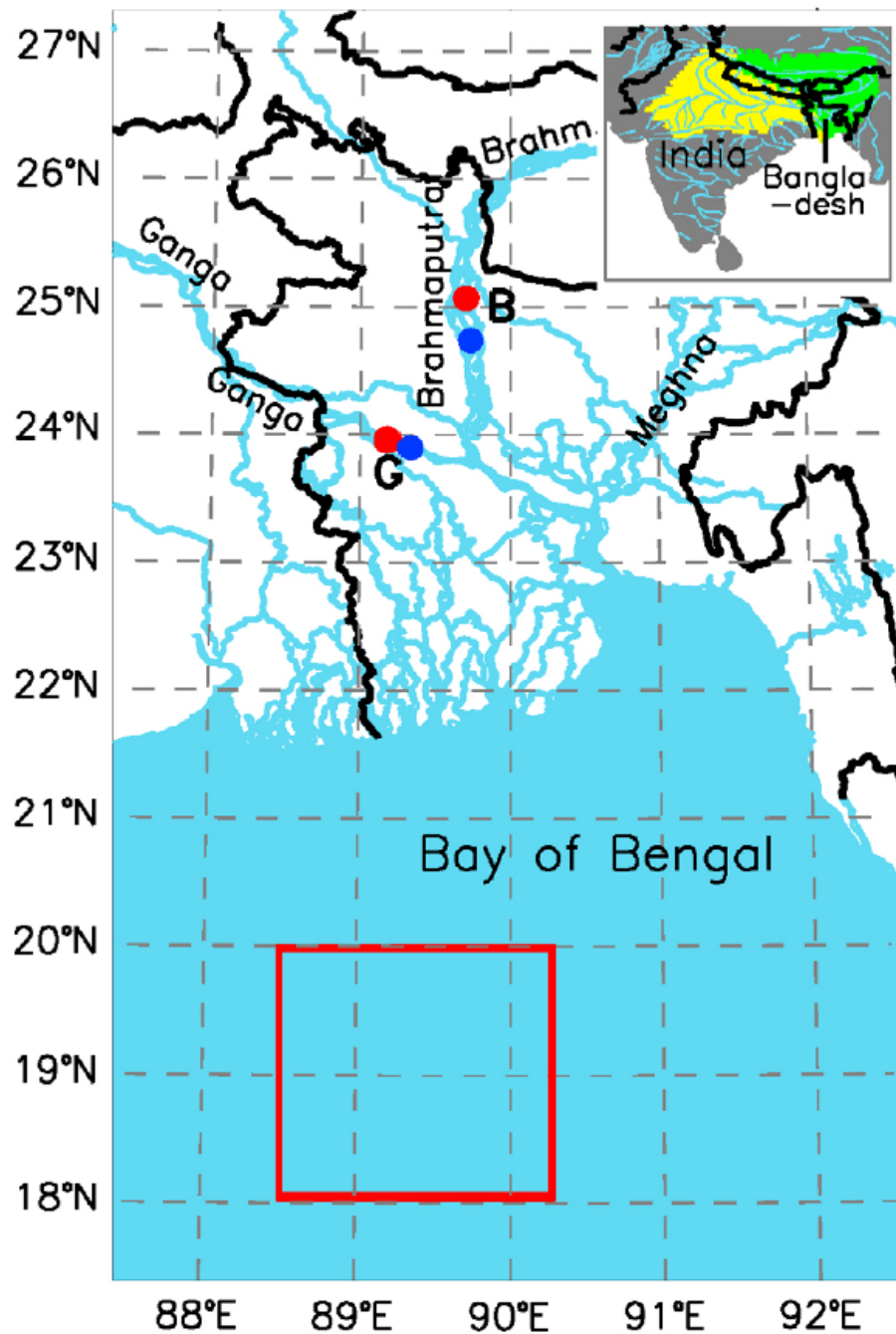
The rainband moves  
north parallel to itself







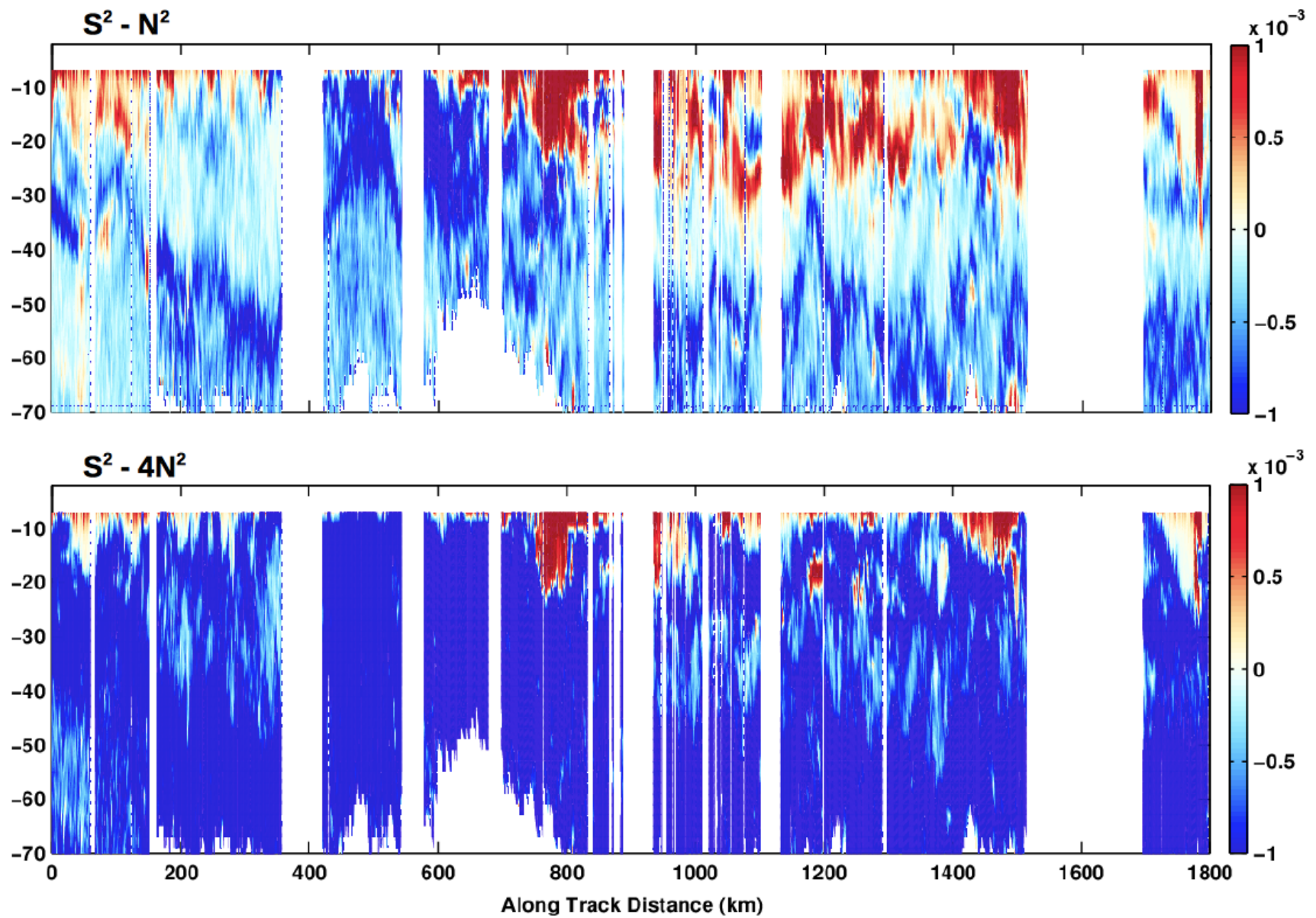




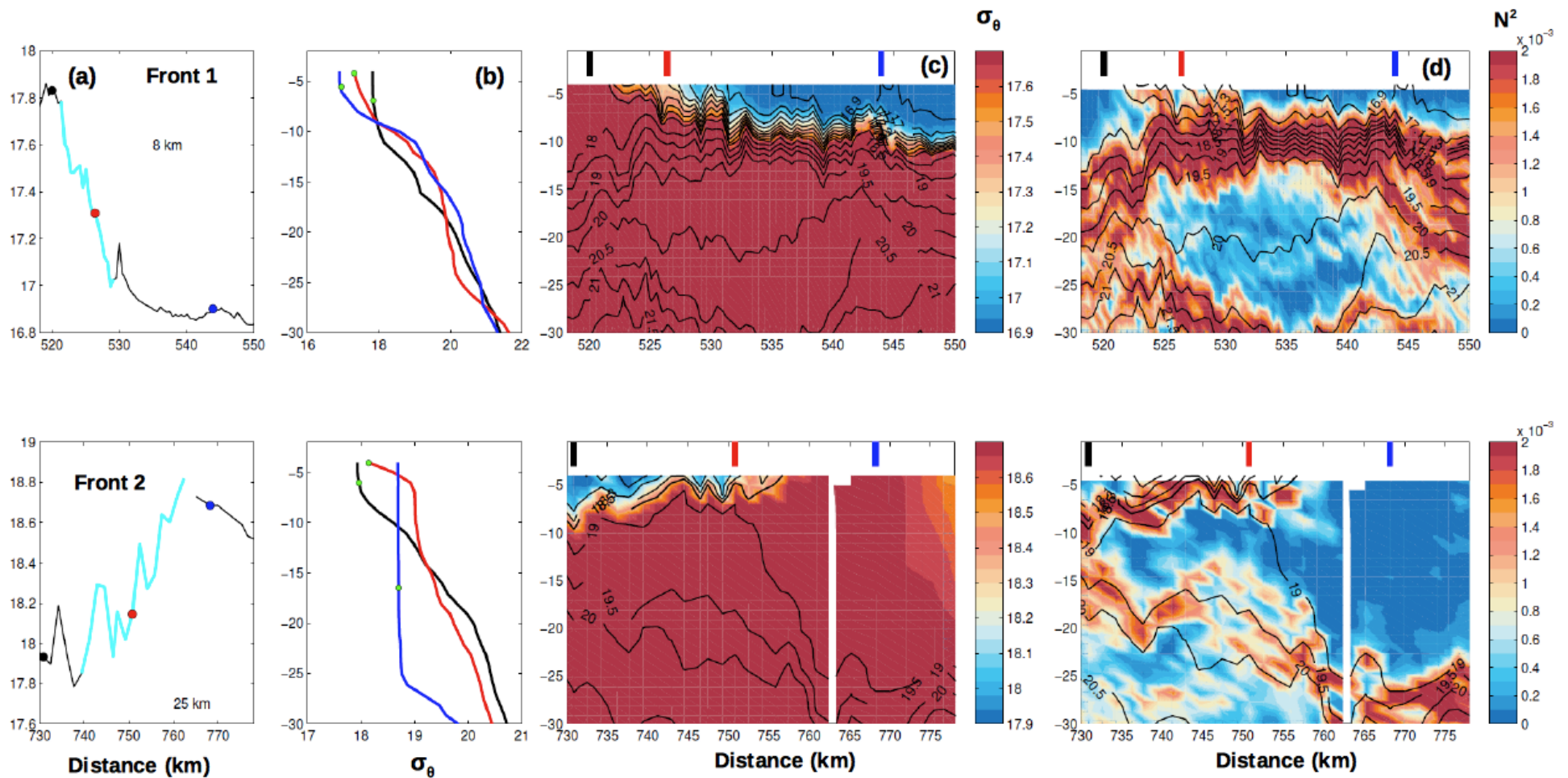
## Rivers

Input of fresh water to the Bay of Bengal  $P-E+R \sim 4000$  cubic km

Papa 2012  
Sengupta 2006



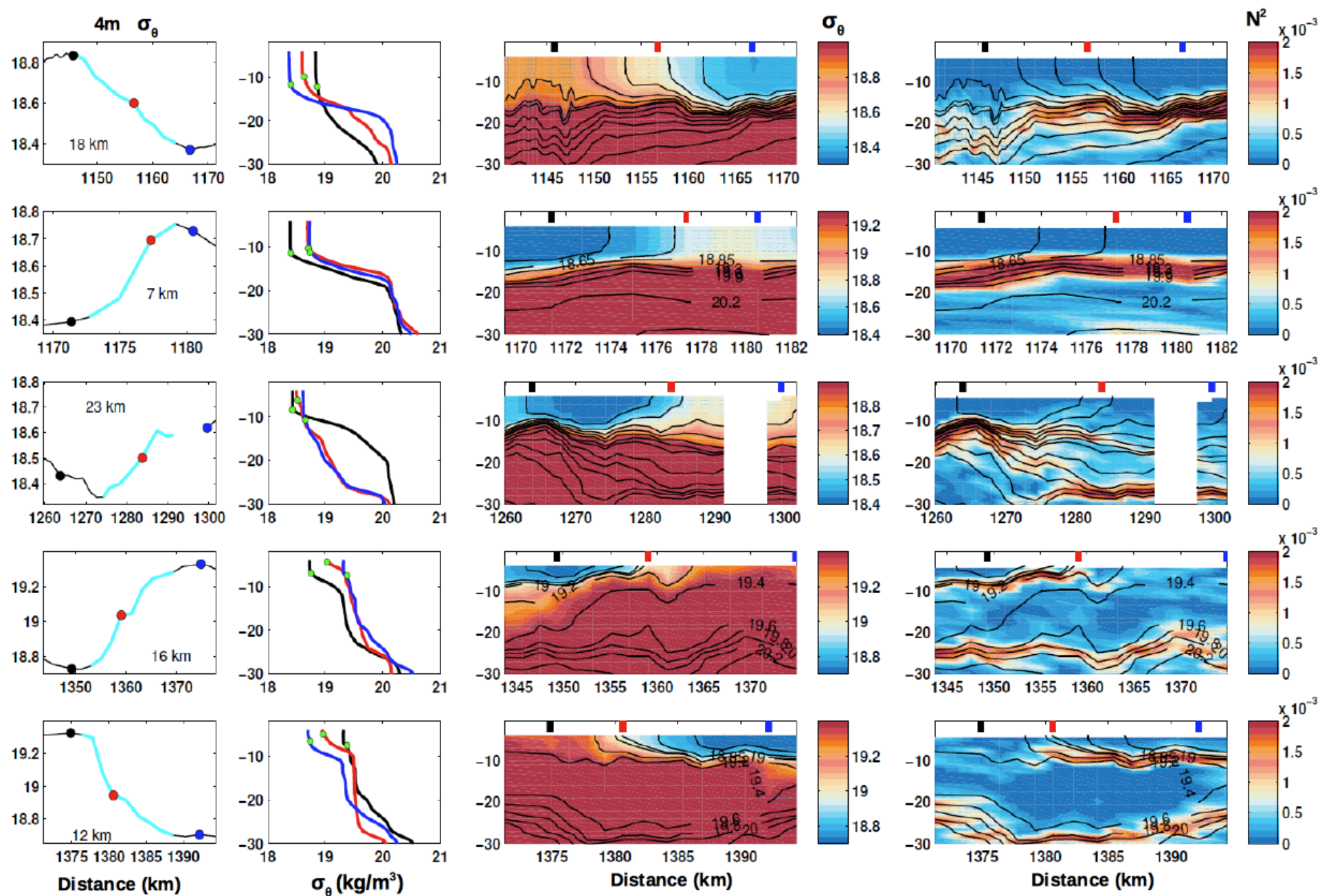
**Reduced shear suggests mixing regions, mostly in the upper 30 m.**



(a) Potential density along the ship track at 4m depth (b)  $\sigma_\theta$  profiles under the front (red) and to the right (blue) and left (black) of the front. Mixed layer depth (MLD) marked by green dot. (c) Depth section of  $\sigma_\theta$  along the ship track, with location of profiles marked on the top. The contours are isopycnals with interval of  $0.5 \text{ kg/m}^3$ . (d)  $N^2$  as function of distance and depth with  $\sigma_\theta$  countours.

### Shallow mixed layers under the fronts

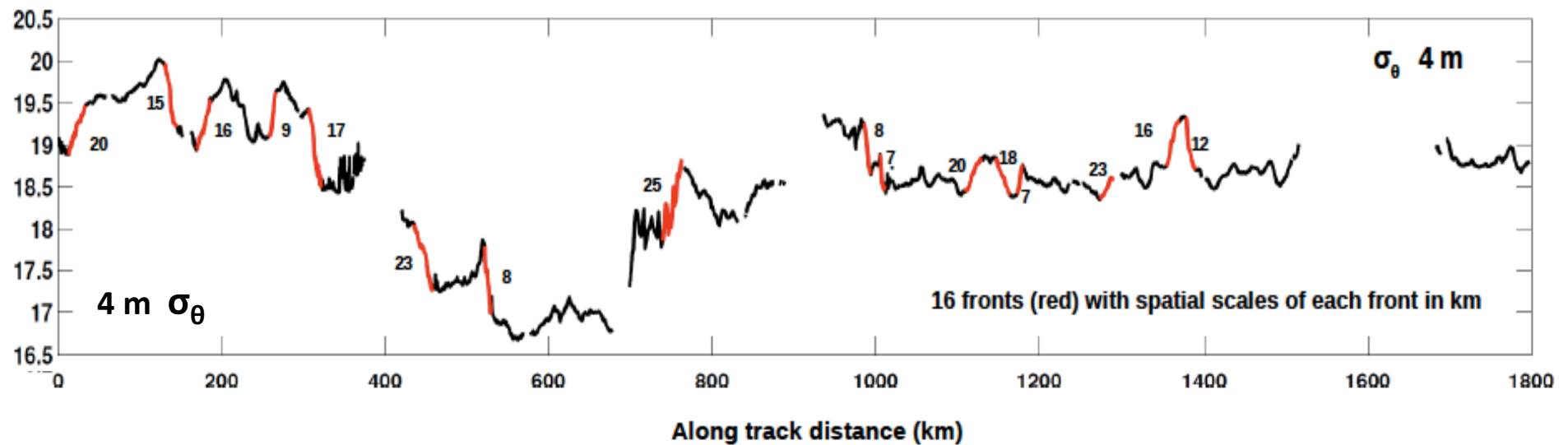




**Shallow mixed layers under 13 out of 16 fronts**

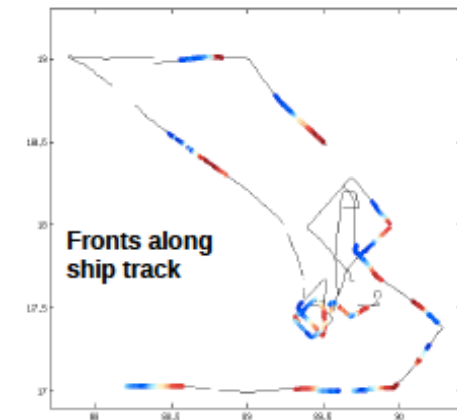


## Salinity dominated density fronts *Sagar Nidhi* SN88



Size ( $\text{kg/m}^3$ )	No of fronts
$0.3 < \text{size} < 0.5$	5
$\text{size} > 0.5$	11

Scale (km)	No of fronts
7 – 20	11
20 - 25	5



Criteria: Total density change across the front (“size”)  $> 0.3 \text{ kg/m}^3$ ,  $|\text{d}\sigma_\theta/\text{dx}| > 1 \text{ std dev}$

16 major fronts, many smaller ones