Exercises using SPEEDY model

CHEN GUANGHUA, JOSHI MANISH KUMAR, SHAHI NAMENDRA KUMAR, VILLEGAS BOLANOS NANCY LILIANA, ZHAO HAIKUN, ZHENG FEI

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Group photo



Outline

• Speedy Model Performance: - Run with

prescribed SSTAs: -

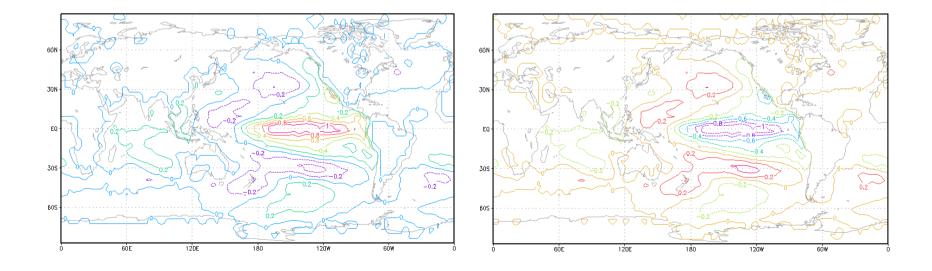
-El Nino & La Nina type.

- TP, Atlantic, & Combined (TP + Atlantic).

A. ENSO teleconnection in SPEEDY model

	SST forcing
NINO	Permanent El Nino
NINA	Permanent La Nina

Simulating period: 1979-2015



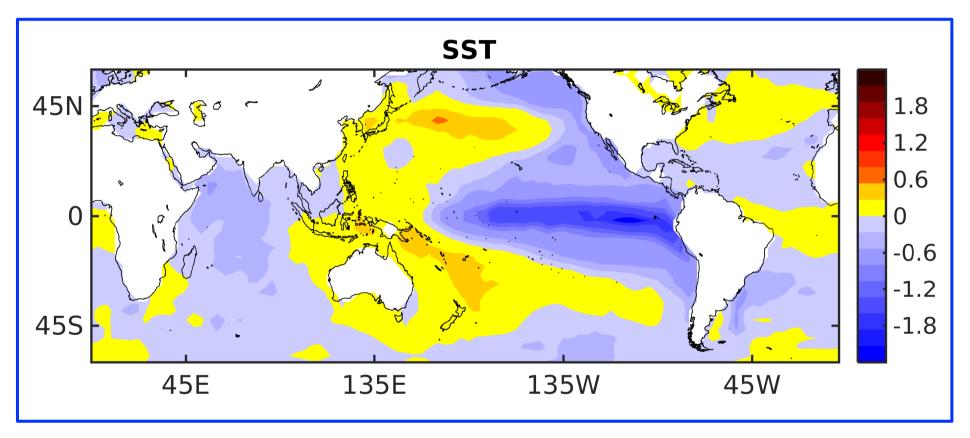


Fig. Climatological Seasonal Mean difference of SST between La Nina and El Nino experiments

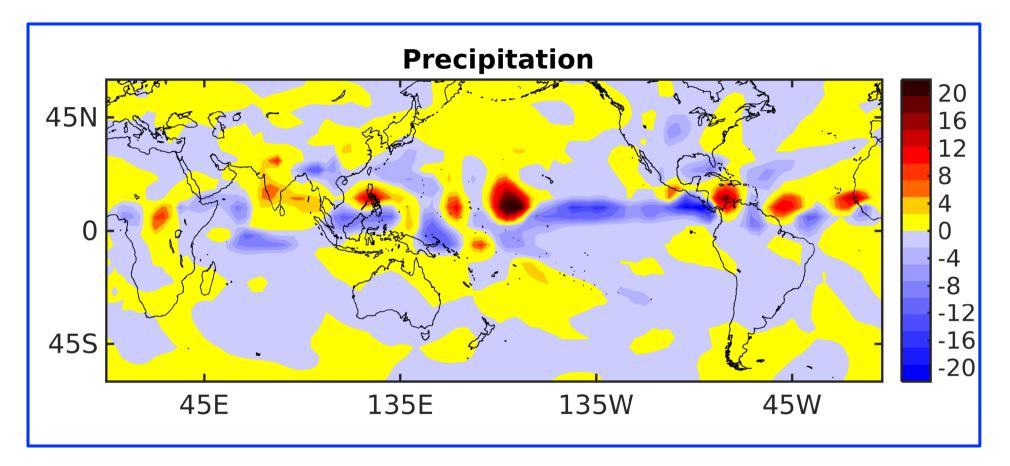


Fig. Climatological Seasonal Mean difference of Precipitation between La Nina and El Nino experiments

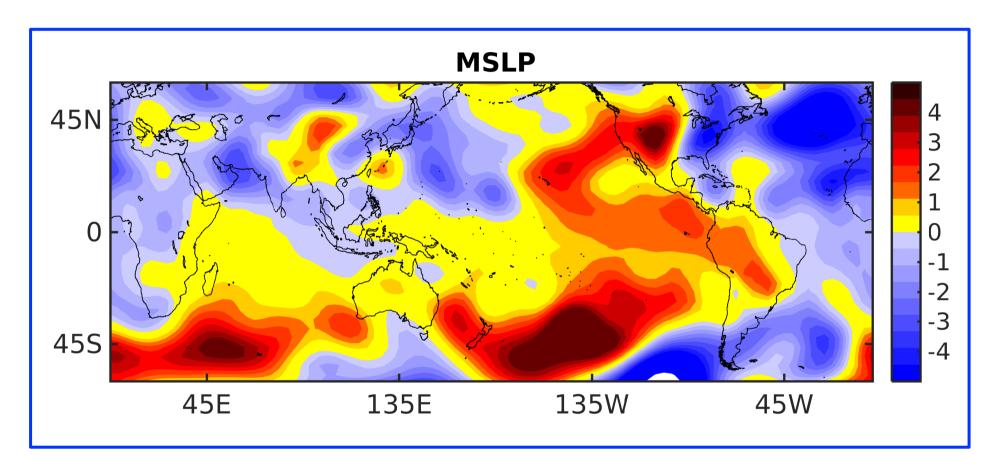


Fig. Climatological Seasonal Mean difference of MSLP between La Nina and El Nino Experiments at lower and upper levels

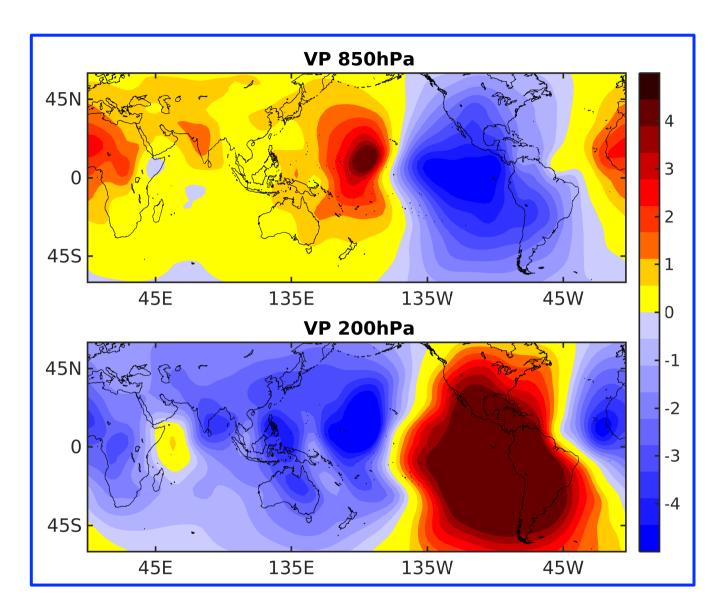
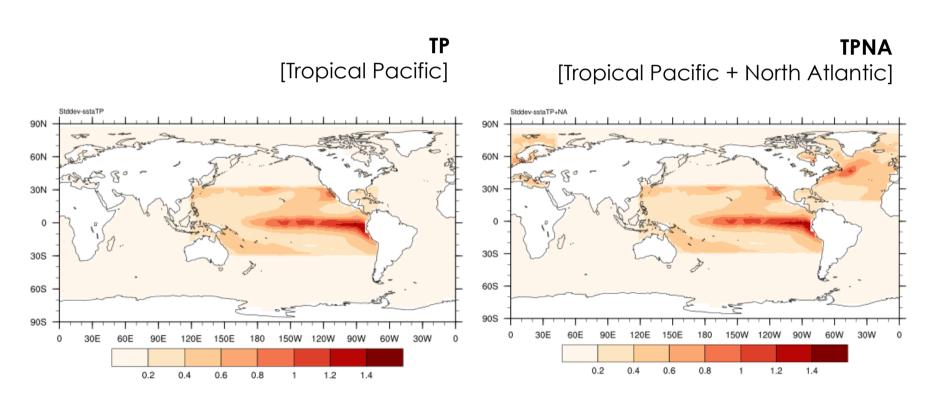


Fig. Climatological Seasonal Mean difference of Velocity Potential between La Nina and El Nino experiments at lower and upper levels

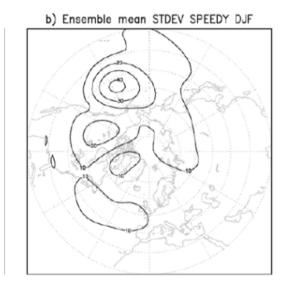
Two types of experiments Each one has 3 ensemble members 1990-2000



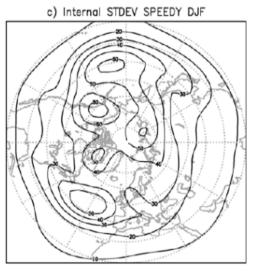
Standard deviation of SST forcing

The variability in SPEEDY AGCM runs can be separated into

A contribution due to the variability of SST forcing



An internal part, due to the intrinsic variability



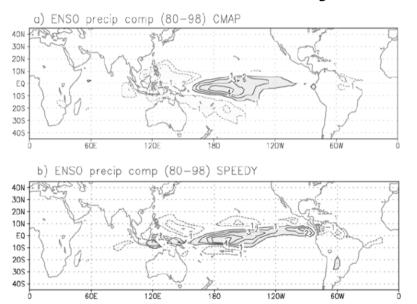
Bracco et al, 2004, Clim. Dyn.

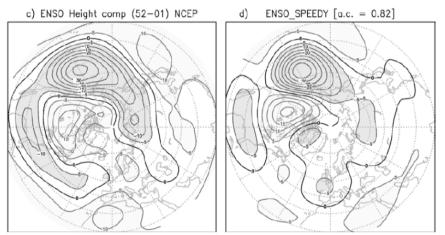
Ensemble mean is used to quantify the response to SST forcing

Internal variability is one source of uncertainty in future climate change (e.g., Hawkins and Sutton 2009; Tebaldi and Knutti 2007; Deser et al. 2012).

B. Role of internal variability in responses of atmospheric circulation to ENSO

The SPEEDY model reproduces well the structure and amplitude of the circulation variability induced by ENSO.





Bracco et al, 2004, Clim. Dyn.

Fig. 2a-d Regression of precipitation and 500 hPa geopotential height onto the NINO3.4 index: a precipitation from CMAP data (period 1988–1998), b precipitation from the SPEEDY ensemble mean (period 1980–1998), c 500 hPa height from NCEP re-analysis data (period 1952–2001), d 500 hPa height from the SPEEDY

ensemble mean (period 1952–2001). In the caption of d the anomaly correlation coefficient of the modelled ENSO composite with c is given in brackets. CI: 1 mm/day for (a) and (b), 5 m for (c) and (d). *Solid lines* (light shading) indicate positive values, negative values are *dotted* (dark shading)

An example of uncertainty in climate projection caused by internal variability

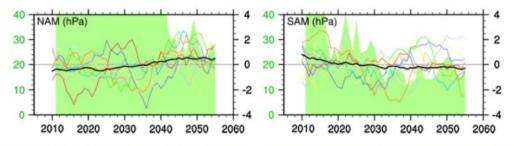
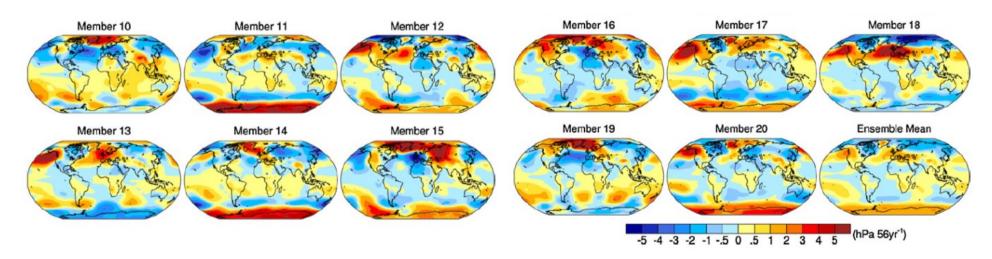


Fig. 5 Ten-year running mean DJF time series of the NAM (*left*) and SAM (*right*), defined as the zonally-averaged SLP anomaly difference between high $(55^{\circ}-90^{\circ})$ and middle latitudes $(30^{\circ}-55^{\circ})$ of the northern and southern hemisphere, respectively. The *thick black*

curve denotes the 40-member ensemble mean, and the *thin colored curves* denote the first 10 ensemble members. The *green shaded curve* shows the minimum number of ensemble members needed to detect a 95% significant change relative to the decade centered on 2010

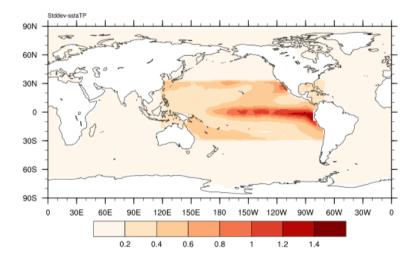
Deser et al. 2012



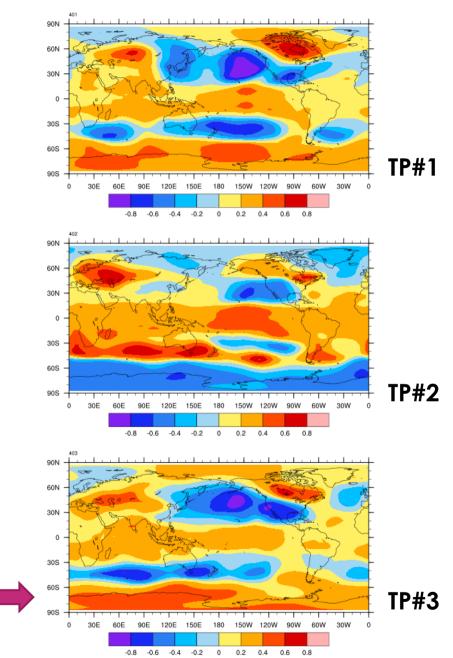
DJF SLP linear trends for individual ensemble members computed over the period 2005-2060

Role of internal variability in the responses of atmospheric circulation to ENSO

[Only using 3 ensemble members in the TP experiment]

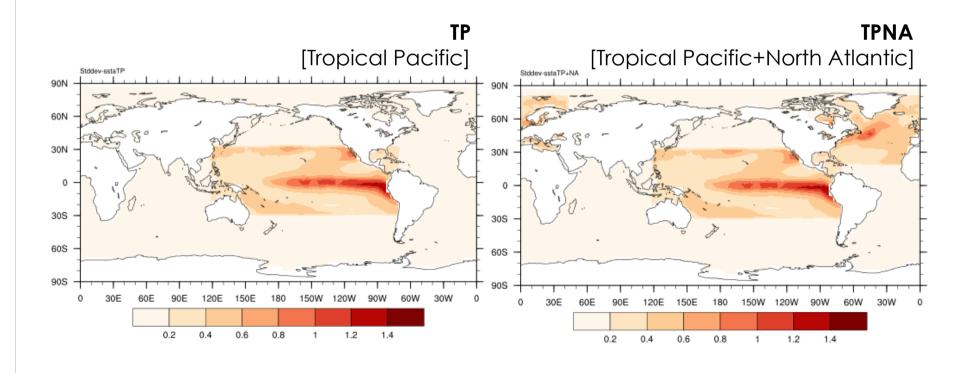


Correlation between the Nino 3.4 index and Z500 in each ensemble member

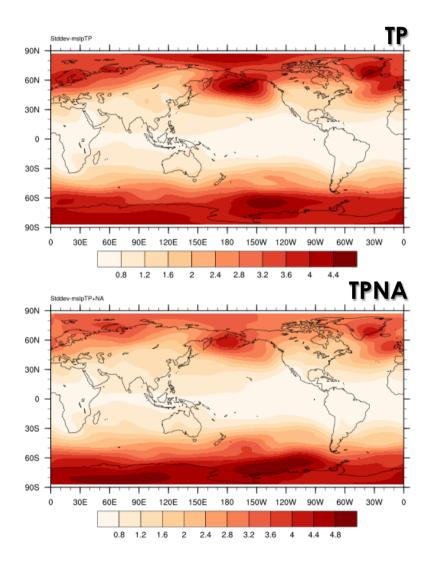


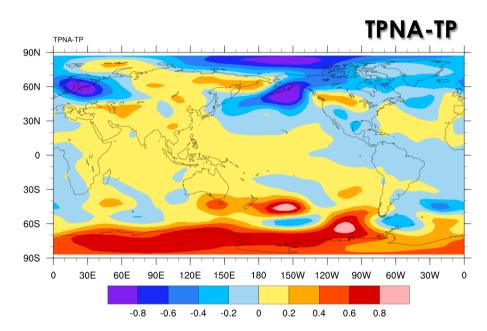
C. Perturbation role of North Atlantic

Compare the ENSO teleconnection in the TP and TPNA experiments

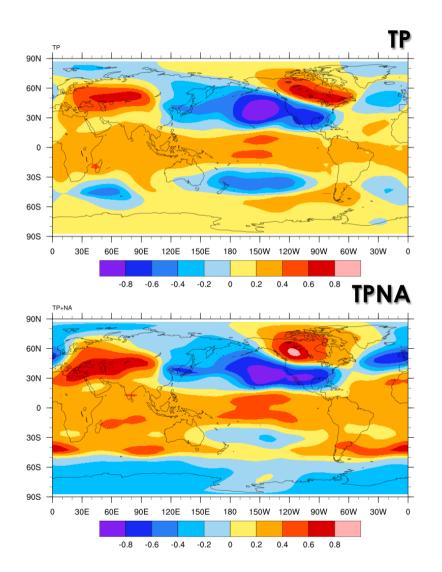


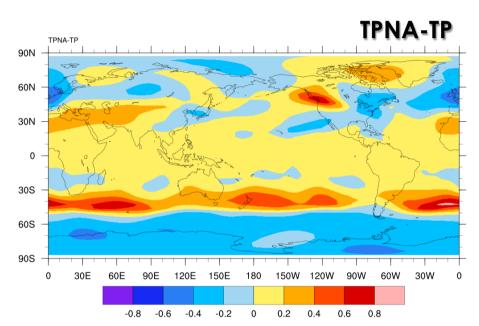
Standard deviation of surface air temperature anomalies [Ensemble mean]



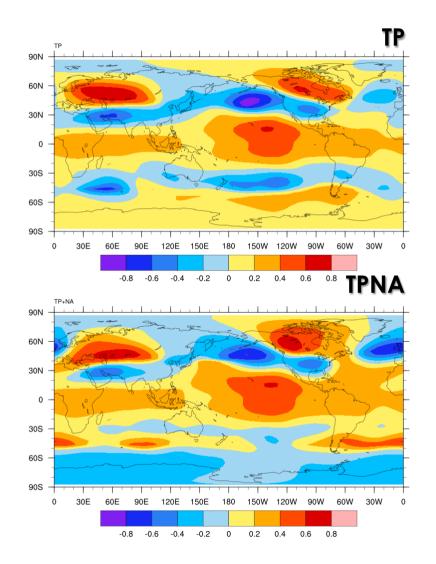


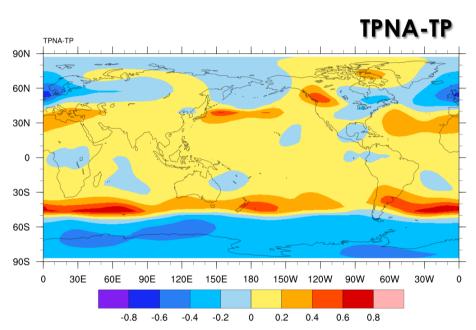
Correlation between the DJF ENSO and DJF Z500 [Ensemble mean]





Correlation between the DJF ENSO and DJF Z200 [Ensemble mean]





Acknowledgements

We'd like to thank all the researchers who help us learn how to use IRI data library and SPEEDY model, and thanks for help us analyze the results.

Thank you!

Correlation between the DJF ENSO and DJF Z850

