Introduction to climate

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Outline

What is the climate?

- 1. The climate: a classification exercise
- 2. The climate: a set of numbers characterizing energy and water fluxes
- 3. The climate: the result of physical laws



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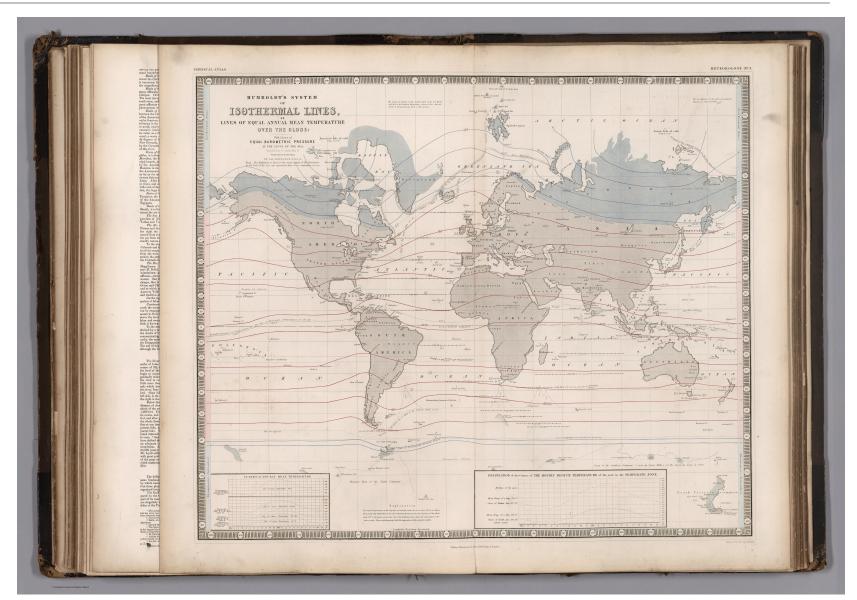


1.1 Alexander von Humboldt classification

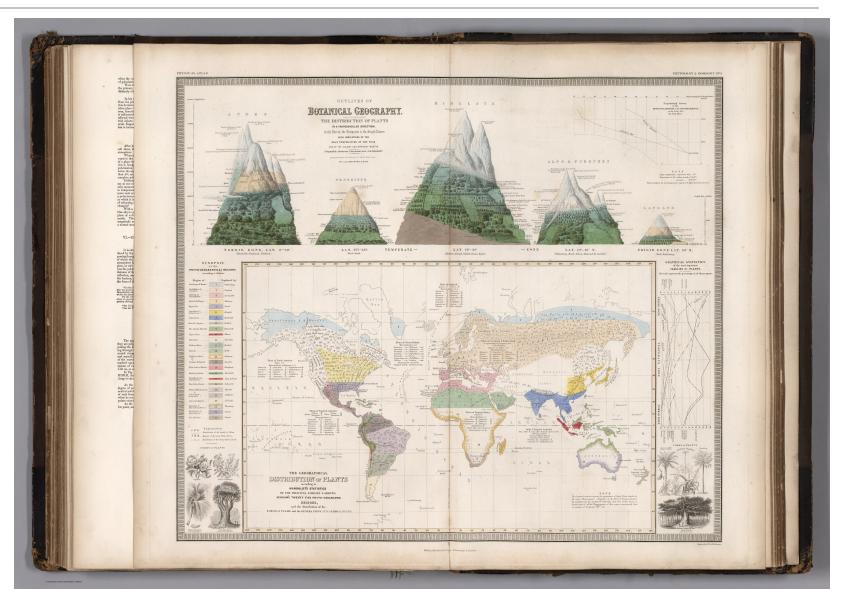
A. v. Humboldt, "Kosmos – Entwurf einer physischen Weltbeschreibung ", 1845-1862



1.1 Alexander von Humboldt classification



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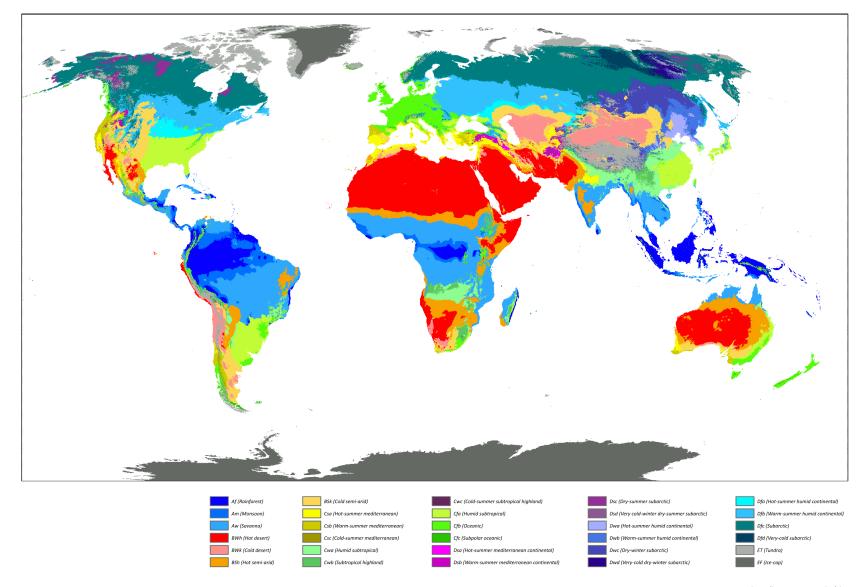
Most well-known climate classification

Köppen classification (1884, 1918, 1936) with later additions by Geiger (1954)

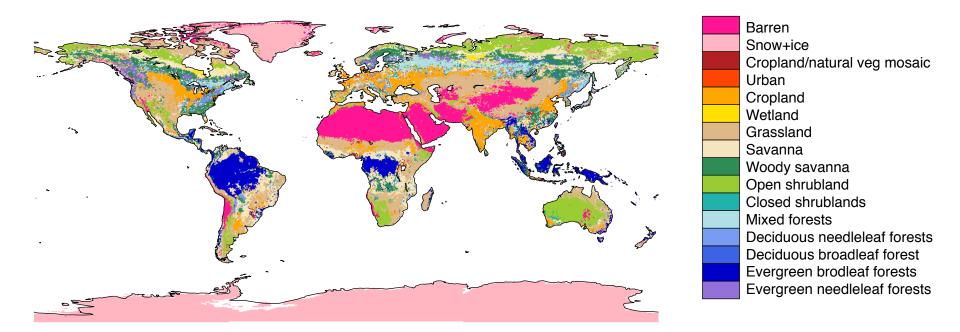


- The Köppen climate classification divides the climate into five main groups
 - Tropical (A), dry (B), temperate (C), continental (D), polar (E)
- Based on monthly averaged temperature (A, C, D, E) or yearly averaged temperature and monthly precipitation (B)
- Each main group is divided into subgroups based on precipitation, e.g. Af, Am, Aw,
 As
- Some of these subgroups are further divided based on temperature, e.g. Cfa, Cfb,
 Cfc











1. The climate: a classification exercise

- Descriptive view
- Based on easily observed properties of the climate system: temperature, precipitation
- Expresses link vegetation-climate
- Indirectly expresses differences in water and energy fluxes
- Doesn't say much about the drivers of the climate system



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	SWd	SWu	LWd	LWu	SH	LH	
TOA	340	100	0	239			_
							in W m ⁻²
SFC	185	25	342	398	21	82	Wild et al. (2015)



	SWd	SWu	LWd	LWu	SH	LH	
TOA	340	100	0	239			
							in W m ⁻²
SFC	185	25	342	398	21	82	Wild et al. (2015)

At equilibrium, energy is conserved:

TOA
$$SWd - SWu + LWd - LWu = 0$$

SFC SWd – SWu + LWd – LWu – SH –LH =
$$0$$

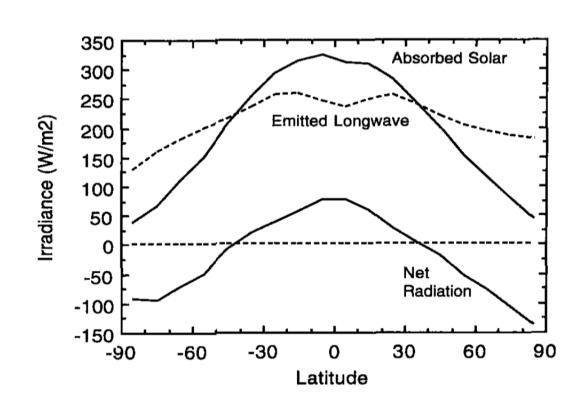
ATM
$$TOA - SFC = 0$$

Atmosphere absorbs 80 W m⁻² solar radiation, emits 183 W m⁻² longwave radiation and absorbs 103 W m⁻² surface fluxes

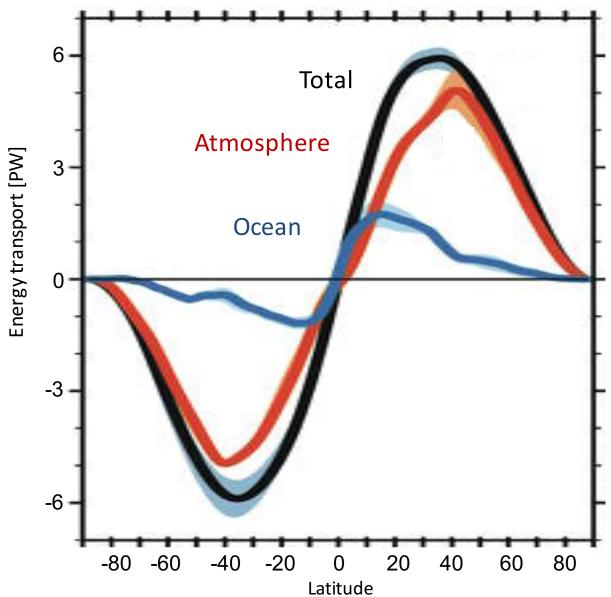


At a point, energy is not balanced ⇒ Storage and Energy transport

- Averaged over a year, the storage term is negligible
- Energy is transported from the tropics to the pole
- Drive atmospheric and oceanic circulation

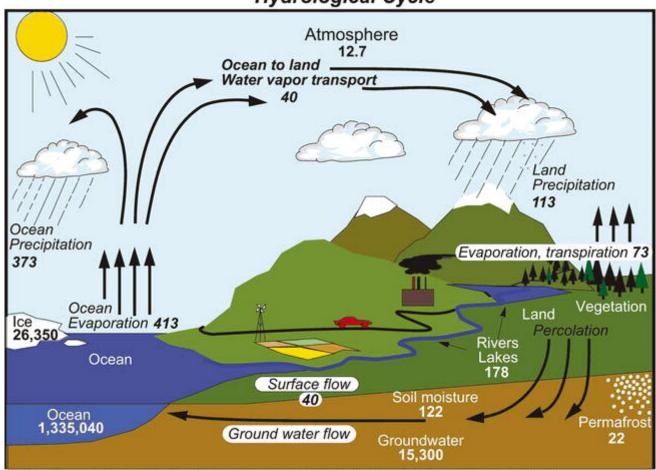




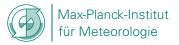




Hydrological Cycle



Units: Thousand cubic km for storage, and thousand cubic km/yr for exchanges



2.2 Water fluxes

Water is also conserved

GLOBAL
$$P = ET$$

Or, using energy budget:

$$P = ET = R_{net}^{sfc} - R_{net}^{toa} - SH$$

 Precipitation approximately compensates radiative cooling

2.2 Water fluxes

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Or, using energy budget:

$$P = ET = R_{net}^{sfc} - R_{net}^{toa} - SH$$

 Precipitation approximately compensates radiative cooling

$$\frac{\partial \mathbf{S}}{\partial t} = \mathbf{P} - \mathbf{ET} - \mathbf{Runoff}$$

ATM

$$\frac{\partial \mathbf{W}}{\partial t} = -\nabla_H \cdot Q - (\mathbf{P} - \mathbf{ET})$$

Q is the vertically integrated twodimensional water vapor flux

2. The climate: a set of numbers

- Descriptive view
- Rely on a combination of satellite, ground measurements and model simulations, still uncertainties on the derived fluxes, in particular at the surface
- Link to underlying conservation of energy and water

TOA SWd – SWu + LWd – LWu = 0 **GLOBAL**
$$P = ET$$

SFC SWd – SWu + LWd – LWu – SH –LH = 0 LAND
$$\frac{\partial S}{\partial t} = P - ET - Runoff$$

ATM TOA – SFC = 0
$$\frac{\partial \mathbf{W}}{\partial t} = -\nabla_H \cdot Q - (\mathbf{P} - \mathbf{ET})$$

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3. The climate: the result of physical laws

How much of the climate properties can we explain by simple conceptual models?

1.	Zero-dim	energy	ba	lance	model	
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2. One-dim energy balance model

3. Radiative equilibrium

4. Radiative convective equilibrium

5. Aquaplanet

Hom. radiation, no atmosphere

Hom. radiation, one-layer atmosphere

Hom .radiation, vertically-resolved atmosphere

Radiation and convection

Radiation, convection, rotation and meridional variations

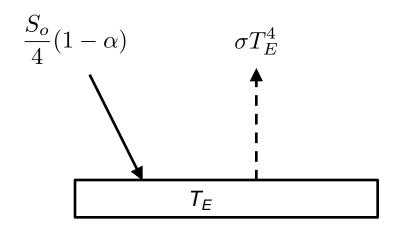


3.1 Zero-dim energy balance model

- Earth (atmosphere+surface) modeled as one layer, no separate atmosphere
- At equilibrium, incoming and outgoing radiative fluxes must compensate:

$$\frac{S_o}{4}(1-\alpha) = \sigma T_E^4$$

$$T_E = 255 \text{ K}$$



- Observed T_F is 288 K!
- In reality, atmosphere makes it difficult for the surface to radiate energy because some of it is absorbed and radiates back down

 S_o solar constant : 1361 W m⁻²

 α albedo : 0.3

 σ Stefan-Boltzmann constant : 5.67 10⁻⁸ W m⁻² K⁻⁴



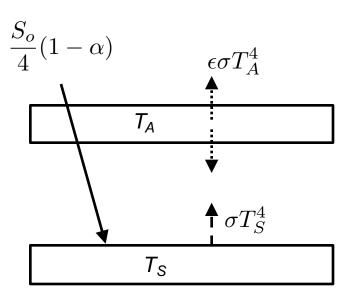
3.2 One-dim energy balance model

- Atmospheric layer separated from the surface
- Atmosphere transparent for shortwave but absorbs and emits longwave, "greenhouse effect"
- Surface radiative balance:

$$\frac{S_o}{4}(1-\alpha) + \epsilon \sigma T_A^4 = \sigma T_S^4$$

Atmospheric radiative balance:

$$\epsilon \sigma T_S^4 = 2\epsilon \sigma T_A^4$$



 S_o solar constant : 1361 W m⁻²

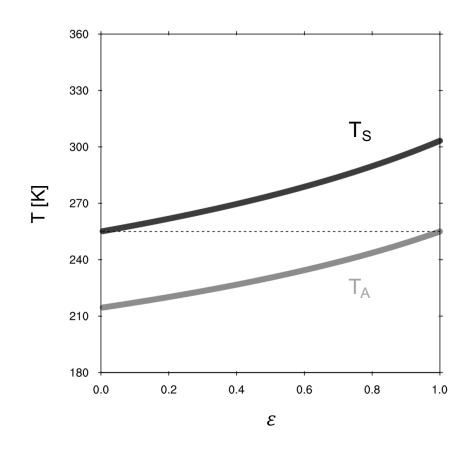
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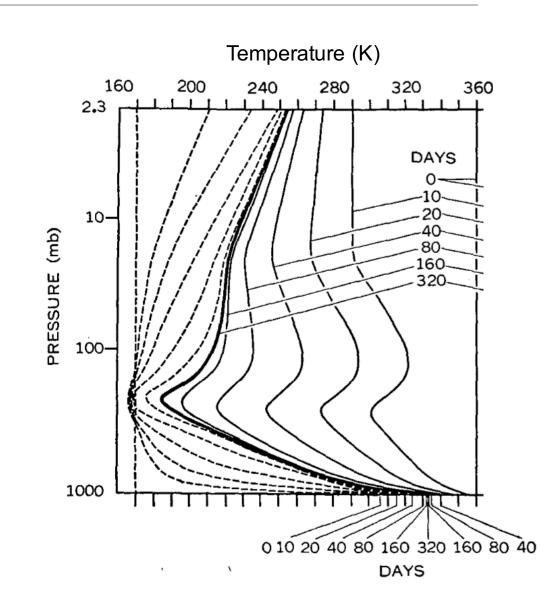
3.2 One-dim energy balance model

 \circ T_s = 288 K for ε =0.77



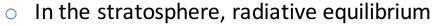
3.3 Radiative equilibrium

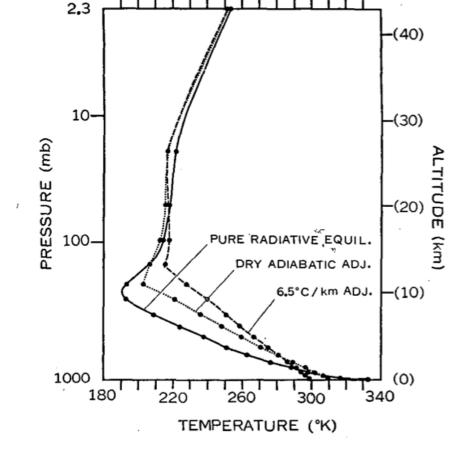
- Add multiple layers in the vertical
- Gaseous absorbers (water vapor and ozone) prescribed
- No cloud
- Annual mean insolation
- Surface albedo of 0.102
- Tropopause too cold
- Surface temperature too warm
- Lapse rate unstable!



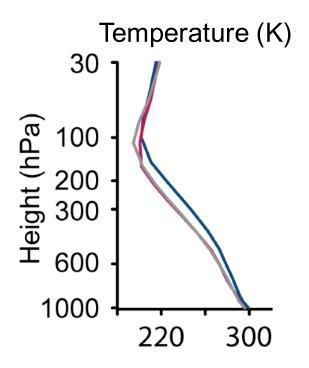


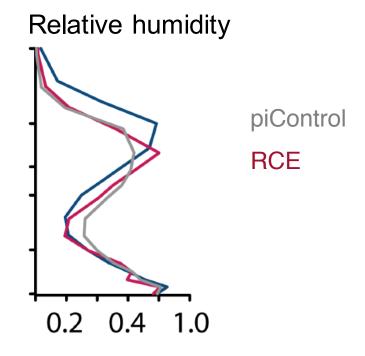
- As soon the temperature profile gets unstable, convection will in reality sets in
- Simple convective parameterization: set lapse rate back to dry or moist adiabat lapse rate
- Able to reproduce the observed mean tropical temperature profile!
- RCE: radiation destabilizes the atmosphere, convection stabilizes it
- Convection is faster than radiation and wins in the troposphere: temperature follows moist adiabat







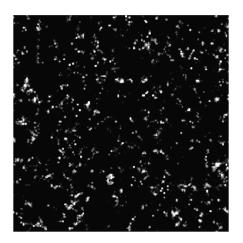






What about the spatial distribution of convection?

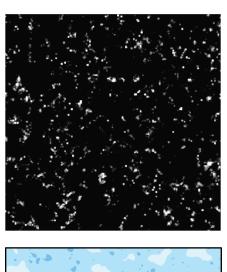
Convection organizes in a blob

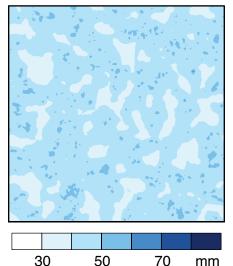




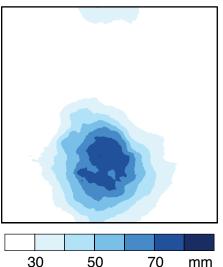
What about the spatial distribution of convection?

- Convection organizes in a blob
- Organization is accompanied by a strong drying





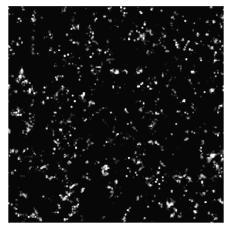


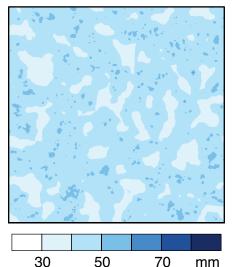


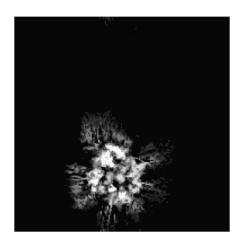


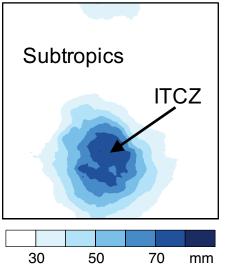
What about the spatial distribution of convection?

- Convection organizes in a blob
- Organization is accompanied by a strong drying
- A simple analog to the ITCZ and its subtropics?



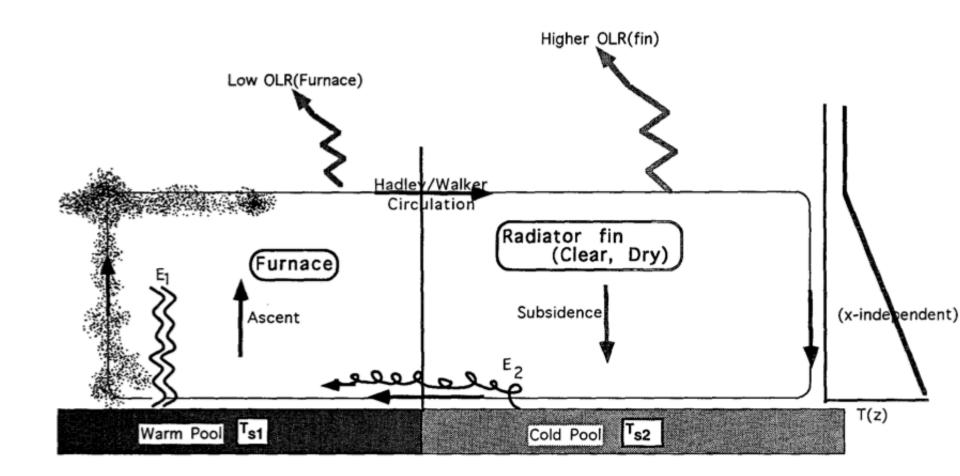








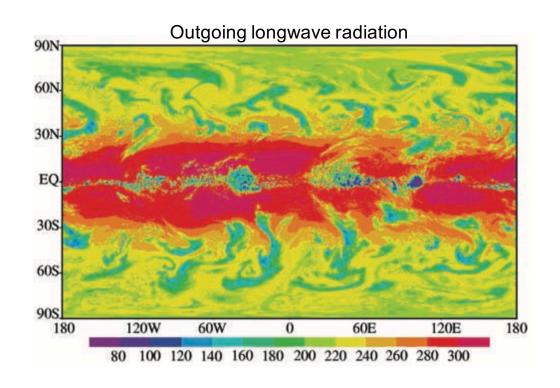
Need the subtropics to avoid falling into a greenhouse state





3.5 Aquaplanet

- Water surface
- Zonally varying SST
- Zonally varying insolation
- Rotation
- Many of the observed features are reproduced





3. The climate: the result of physical laws

- Simple conceptual model can reproduce basic properties of the climate system
- Observed surface temperature can be understood from a 1-dimensional energy balance model, but needs greenhouse effect
- Observed stratospheric tropical temperature profile can be understood from pure radiative equilibrium
- Observed tropospheric tropical temperature profile can be understood from a statistical equilibrium between convection and radiation
- But need dry areas to cool the planet
- Need aquaplanet set-up for extratropical storms to develop



Conclusions

What is the climate?

- 1. The climate: a classification exercise
 - Descriptive view based on easily observed quantities whose variations result from energy/water fluxes
- 2. The climate: a set of numbers characterizing energy and water fluxes
 - Energy and water balances provide constraints on the climate
- The climate: the result of physical laws
 - Greenhouse effect, Radiative Convective Equilibrium, Furnaceradiator fin, aquaplanet

