

Super-parameterization: what it is and what is “super” about it?

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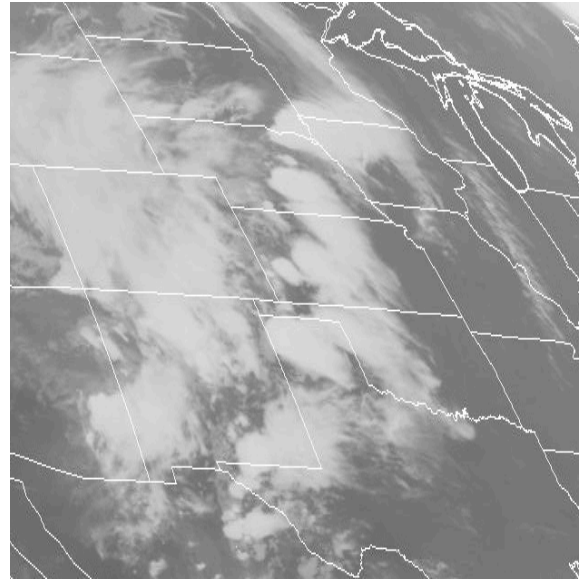
Boulder, Colorado, USA



This material is based upon work supported by the National Center for Atmospheric Research, which is a major facility sponsored by the National Science Foundation under Cooperative Agreement No. 1852977.

- Introduction: the concept of super-parameterization (SP)
- Examples of initial applications
- Further developments and applications
- Towards global LES: can we get there faster?

**Mesoscale convective
systems over US**

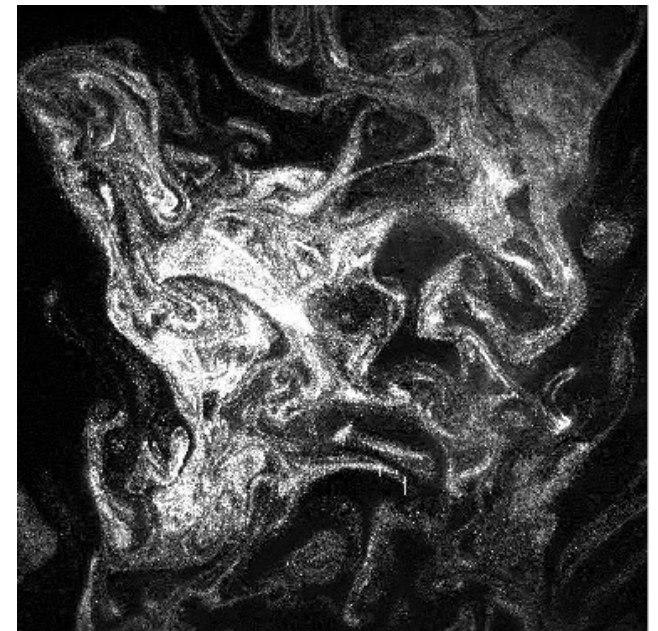


**Mixing in laboratory
cloud chamber**

*Clouds and climate:
the range of scales...*



**Small cumulus
clouds**

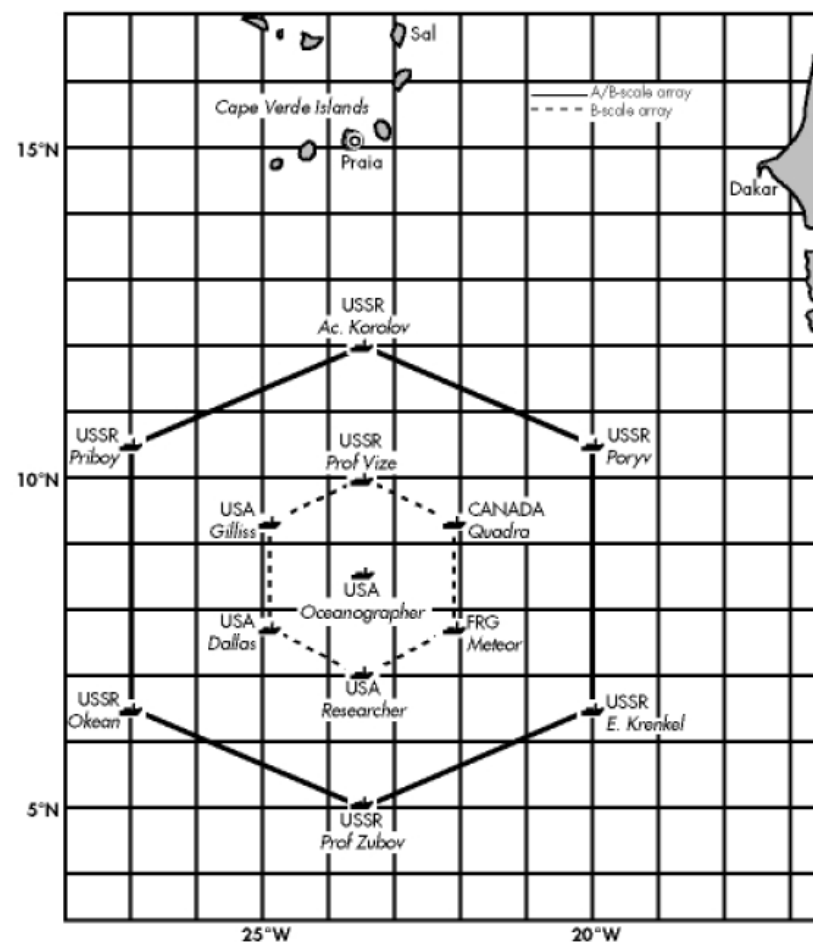
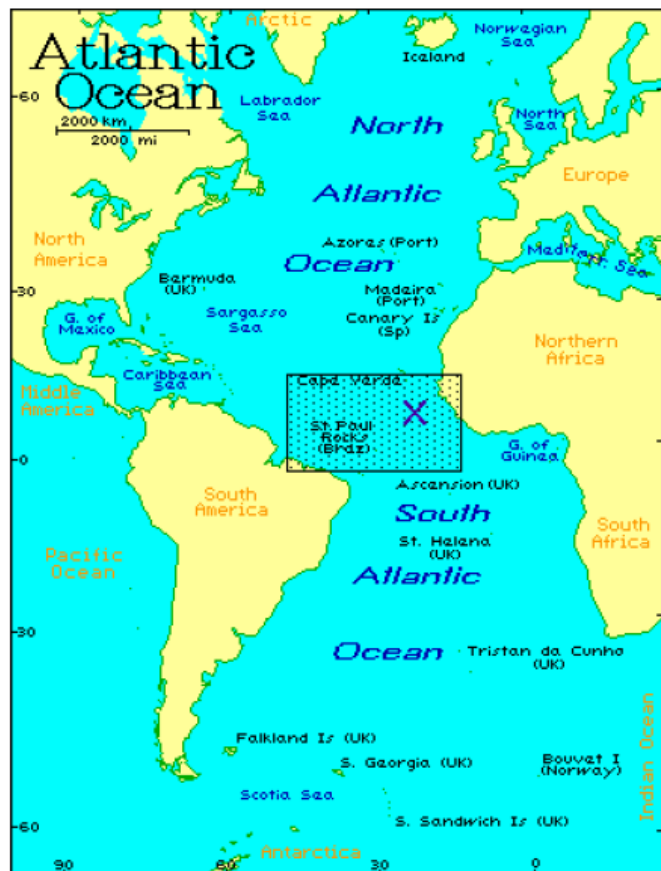


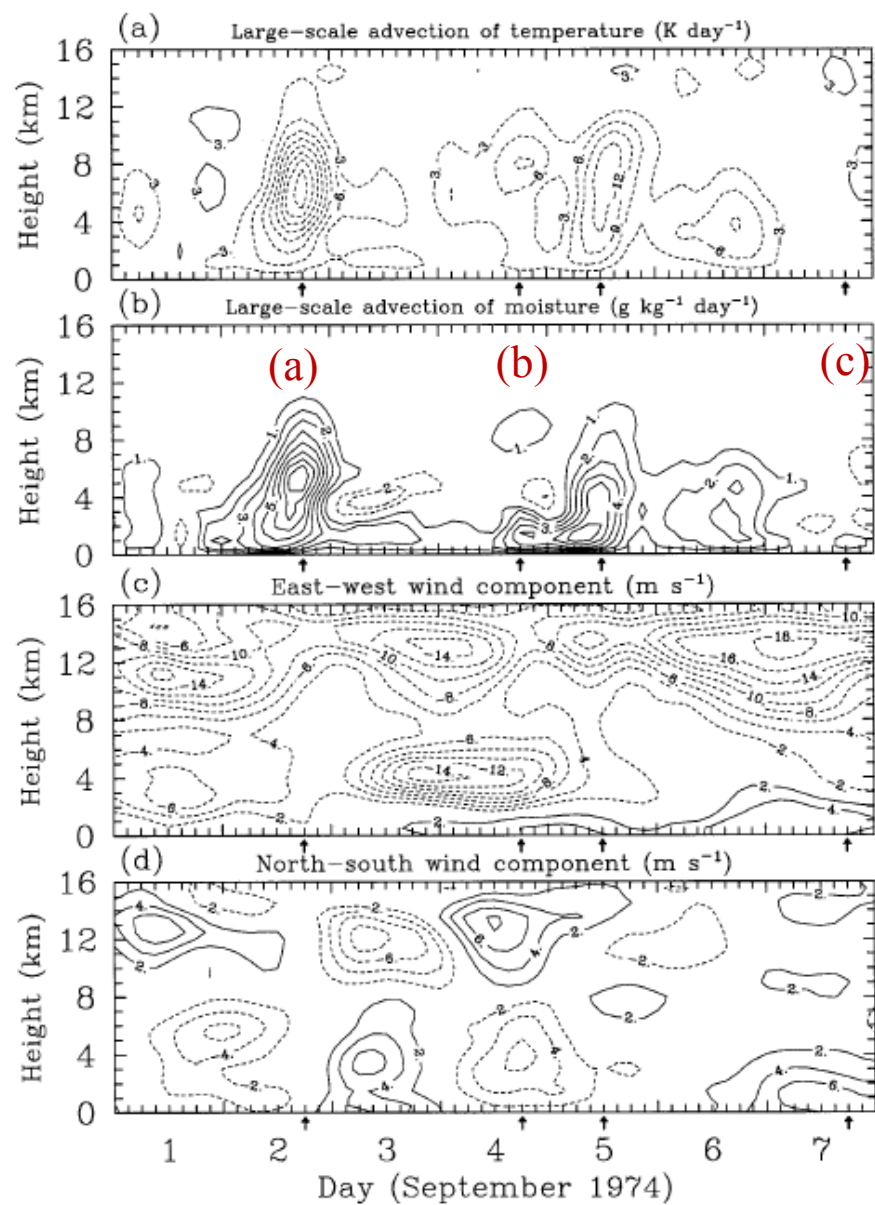
10 cm



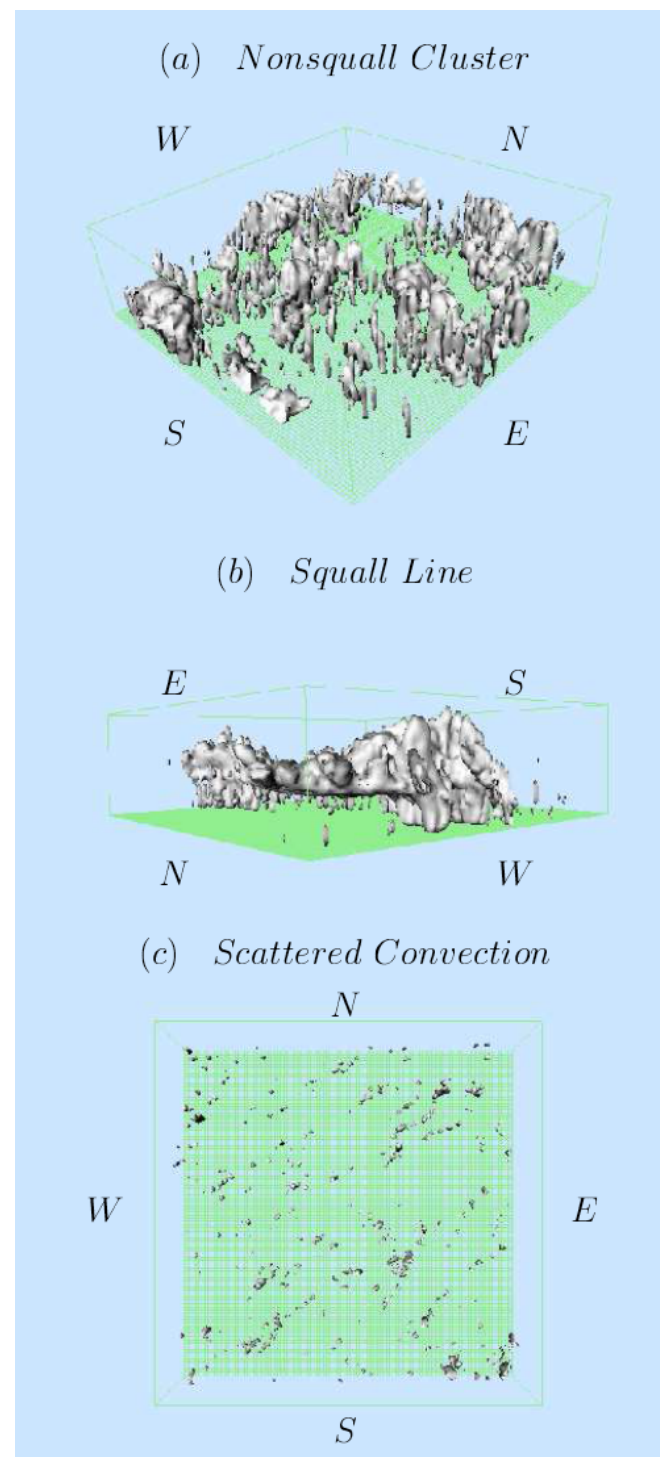
June 15, 1974 to September 23, 1974

Project Location: Atlantic Ocean





Grabowski et al. *JAS* 1996, 1998a,b



Cloud-resolving modeling of GATE cloud systems (Grabowski et al. JAS 1996, 1998)

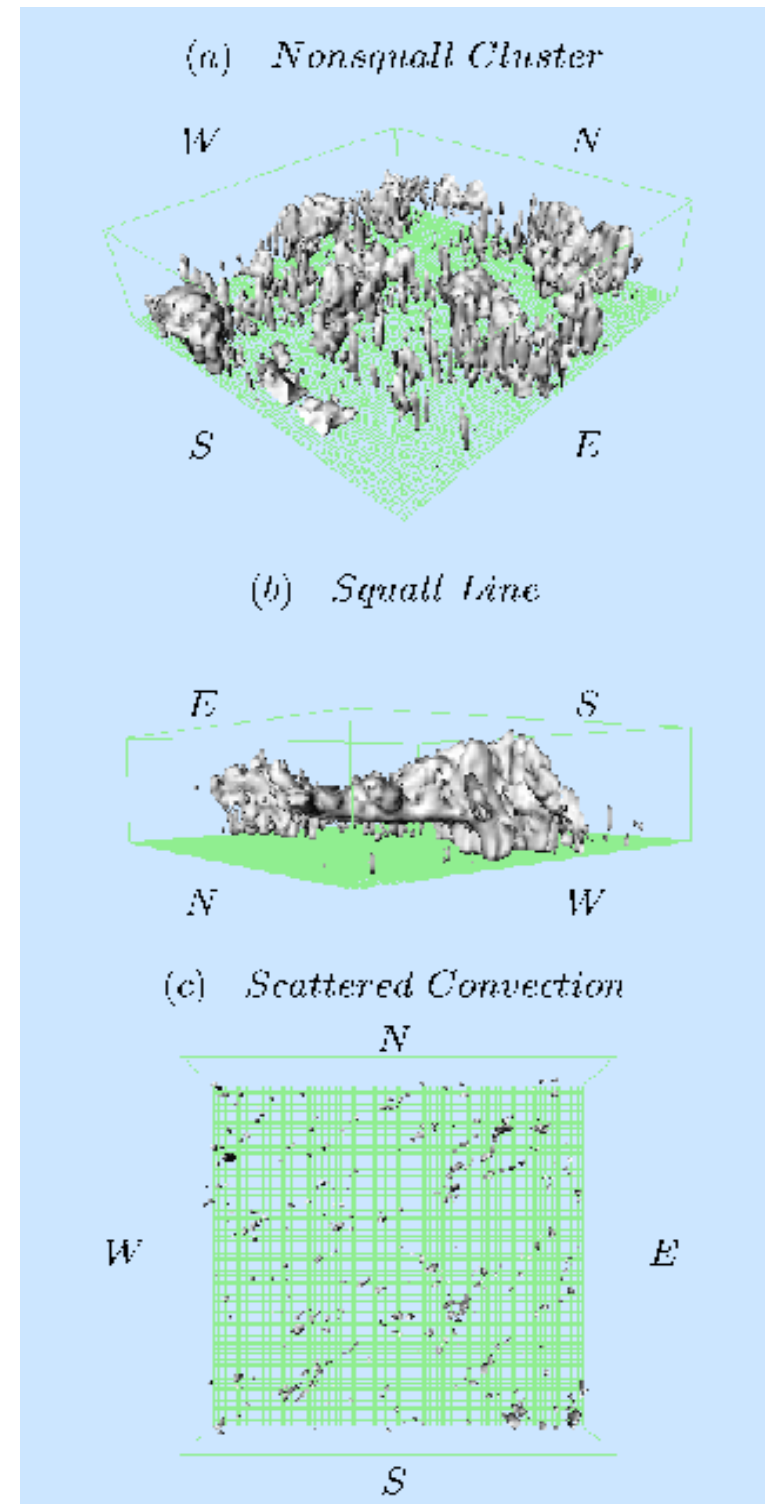
2 Sept, 1800 Z

400 x 400 km horizontal domain,
doubly-periodic,
2 km horizontal grid length

Driven by observed large-scale conditions

4 Sept, 1800 Z

7 Sept, 1800 Z



Grabowski et al. JAS 1998:

“...low resolution two-dimensional simulations can be used as realizations of tropical cloud systems in the climate problem and for improving and/or testing cloud parameterizations for large-scale models...”

- *Can we use 2D cloud-resolving model (CRM) in all columns of a climate model to represent deep convection?*
- *Can we move other parameterizations (radiative transfer, land surface model, etc) into 2D CRM?*

Cloud-Resolving Convection Parameterization (CRCP) (super-parameterization, SP)

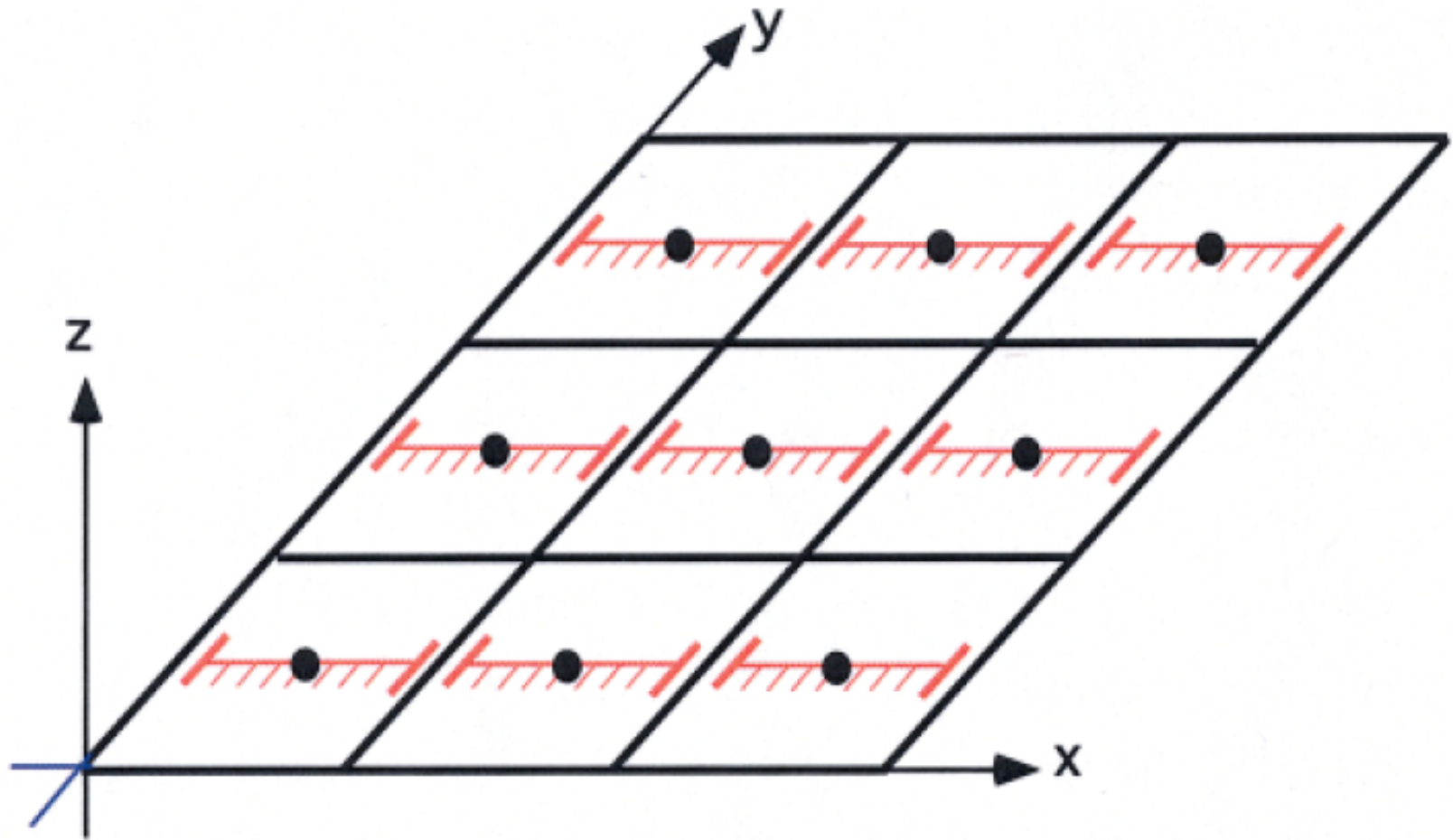
Grabowski and Smolarkiewicz, Physica D 1999

Grabowski, JAS 2001

The idea is to represent subgrid scales of the 3D large-scale model (horizontal resolution of 100s km) by embedding periodic-domain 2D CRM (horizontal resolution around 1 km) in each column of the large-scale model

Another (better?) way to think about CRCP: CRCP involves hundreds or thousands of 2D CRMs interacting in a manner dictated by the large-scale dynamics

Original SP proposal:



- **CRCP is a “parameterization”** because scale separation between large-scale dynamics and cloud-scale processes is assumed; cloud models have periodic horizontal domains and they communicate only through large scales
- **CRCP is “embarrassingly parallel”**: a climate model with CRCP can run efficiently on 1000s of processors
- **CRCP is a physics coupler**: most (if not all) of physical (and chemical, biological, etc.) processes that are parameterized in the climate model can be included into CRCP framework

“A day, a year, a millennium” paradigm

With the same amount of computer time, one can perform:

- about a day-long simulation using cloud-resolving AGCM
- about a year-long climate simulation using AGCM with super-parameterization
- about a millennium-long climate simulation using a traditional AGCM with parameterized convection

CRCP (SP, MMF) was making a steady progress...

- **Grabowski (NCAR): idealized simulations of large-scale tropical dynamics (MJO; Grabowski JAS 2003, 2006; Grabowski and Moncrieff QJ 2004)**
- **Khairoutdinov/Randall (CSU): realistic climate simulations using CAM (atmospheric part of NCAR's CCSM; Khairoutdinov et al. JAS 2005, 2007)**
- **Arakawa: proposal to extend original formulation to remove some of the limitations (see Randall et al. BAMS 2003, Jung and Arakawa MWR 2005)**
- **Effort within ARM Program to compare SP AGCM simulations with ARM observations (CSU model in DOE Labs, e.g., Ovtchinnikov et al. JCli 2006)**
- **Efforts within NASA (Goddard, Langley) to run SP GCMs (Tao, Xu)**
- **NSF Science and Technology Center at CSU: Center for Multiscale Modeling of Atmospheric Processes, CMMAP**



CMMAP

Center for Multi-Scale Modeling of Atmospheric Processes

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August Team Meeting

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A Multiscale Model Is A Scientific Water Hole.



(Dave Randall, 2007)

Examples of initial applications:

- Simulations of the Madden-Julian Oscillation (MJO)-like coherences on a constant-SST aquaplanet (Grabowski JAS 2001, 2006)
- AGCM simulations using CAM (Colorado State University: Khairoutdinov et al. JAS 2005; JCLI 2007)

Madden and Julian, JAS 1972

34 days

0 days, 45 days

12 days

22 days

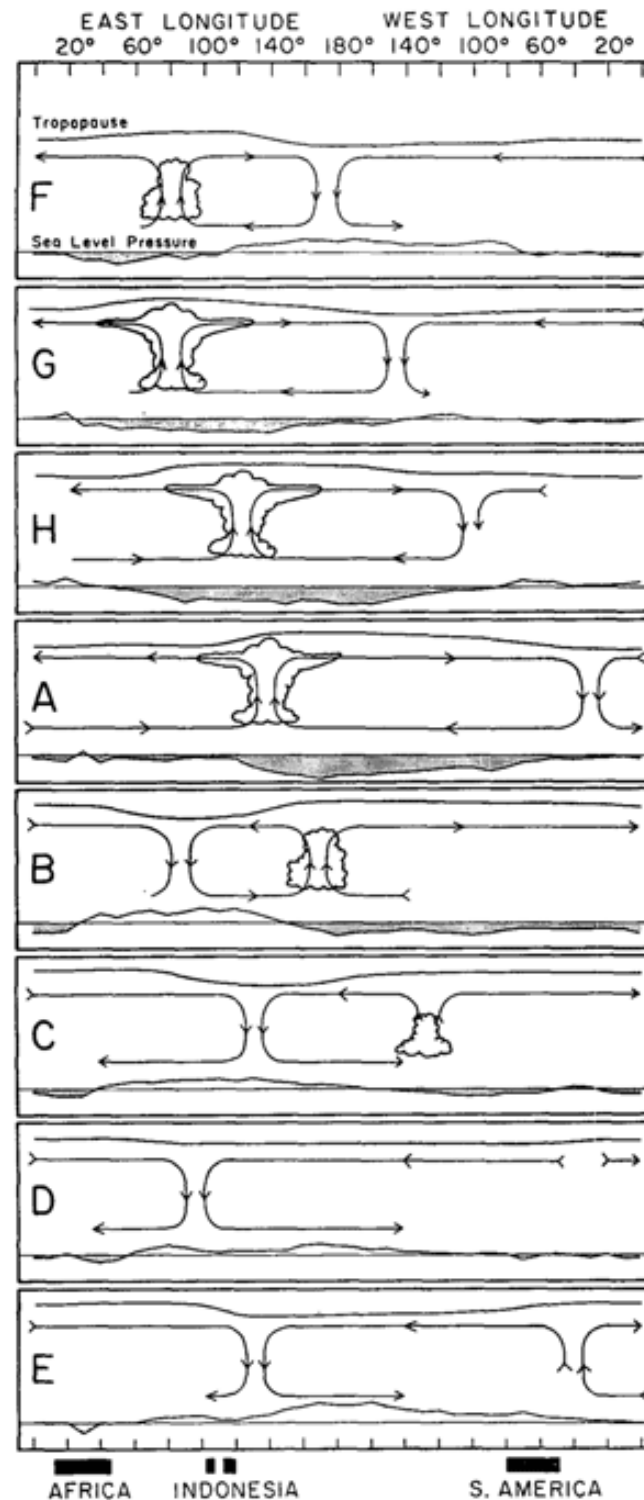
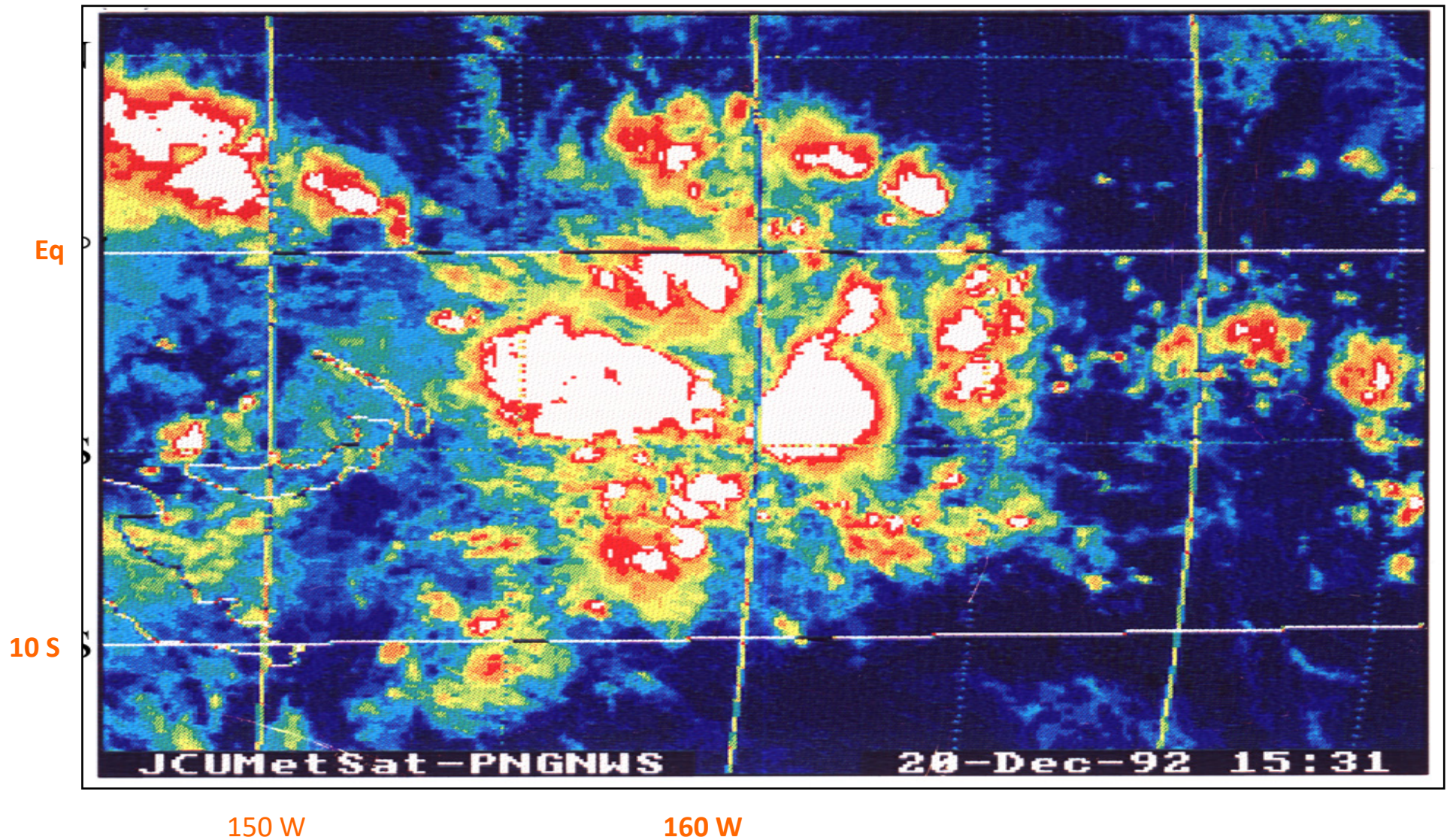
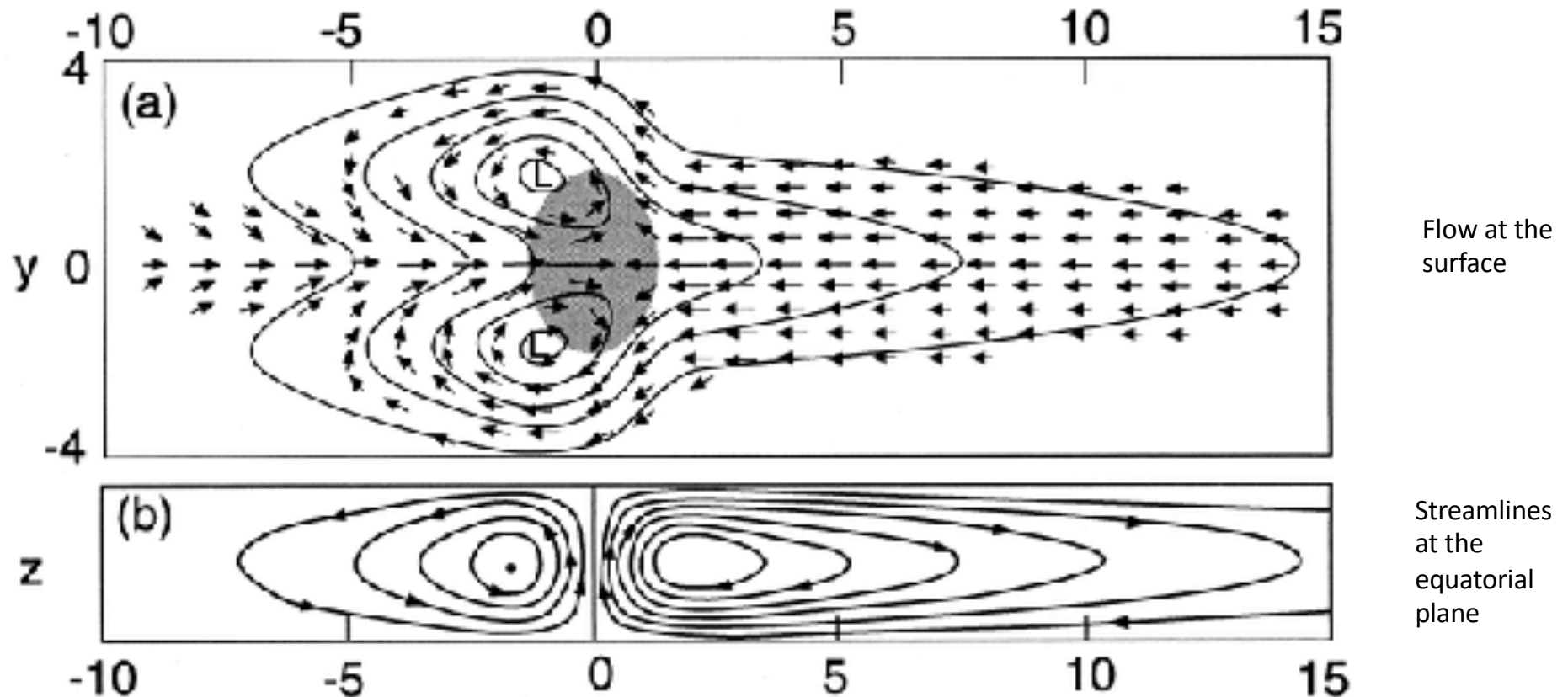


FIG. 16. Schematic depiction of the time and space (zonal plane) variations of the disturbance associated with the 40-50 day oscillation. Dates are indicated symbolically by the letters at the left of each chart and correspond to dates associated with the oscillation in Canton's station pressure indicated in Fig. 11. The mean pressure disturbance taken from Fig. 12 is plotted at the bottom of each chart with negative anomalies shaded. The circulation cells are based on the mean zonal wind disturbance presented in Fig. 13. Regions of enhanced large-scale convection are indicated schematically by the cumulus and cumulonimbus clouds. The relative tropopause height is indicated at the top of each chart.

Satellite picture of a super-cluster during TOGA COARE





Circulation produced by deep heating anomaly over the equator (stripped), with Kelvin-wave response to the east and Rossby-wave response to the west, the Kelvin-Rossby wave (Gill 1980, as shown by Salby 1996).

Plethora of theories trying to explain the large-scale organization of tropical convection:

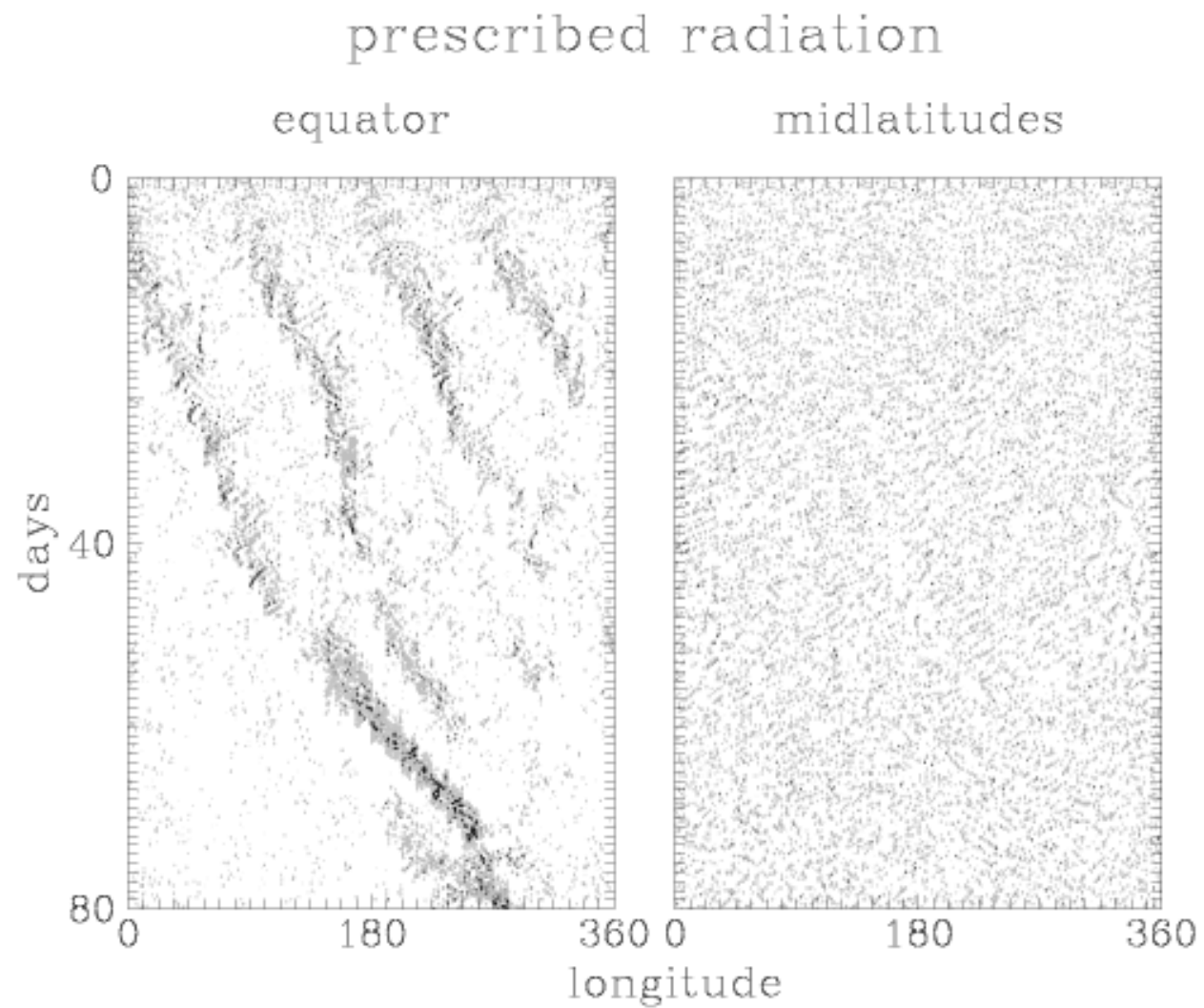
- Coupling between convection and large-scale equatorial perturbations (wave-CISK, etc; e.g., Lindzen 1974; Lau et al. 1989; Wang and Rui 1980; Majda and Shefter 2001...)
- Impact of moisture/clouds on radiative transfer (e.g., Pierrehumbert 1995; Raymond 2000, 2001...)
- Impact of free-tropospheric humidity on convection (e.g., Raymond 2000; Tompkins 2001a,b; Grabowski 2003; Grabowski and Moncrieff 2004; Bony and Emanuel 2005)
- Impact of gravity waves on subsequent convective development (e.g., Mapes 1993, 1998; Ouchi 1999)
- Up-scale effects of organized convection (Moncrieff 2004) and synoptic-scale waves (Biello and Majda 2005)
- Atmosphere-ocean interaction:
 - WISHE (Emanuel 1997; Neelin et al. 1997)
 - coupled atmosphere-ocean dynamics (e.g., Flatau et al. 1997)

MJO-like coherent structures on a constant-SST (“tropics everywhere”) aquaplanet

- Size and rotation as Earth
- SST=30 degC
- Prescribed radiative cooling or interactive radiation transfer model (within CRCP domains; sun overhead over entire aquaplanet, no diurnal cycle)
- Atmosphere at rest (at large scales) at $t=0$
- Low horizontal resolution global model (32 x 16), small cloud models (100 x 50; $dx=2$ km, $dz=0.5$ km)

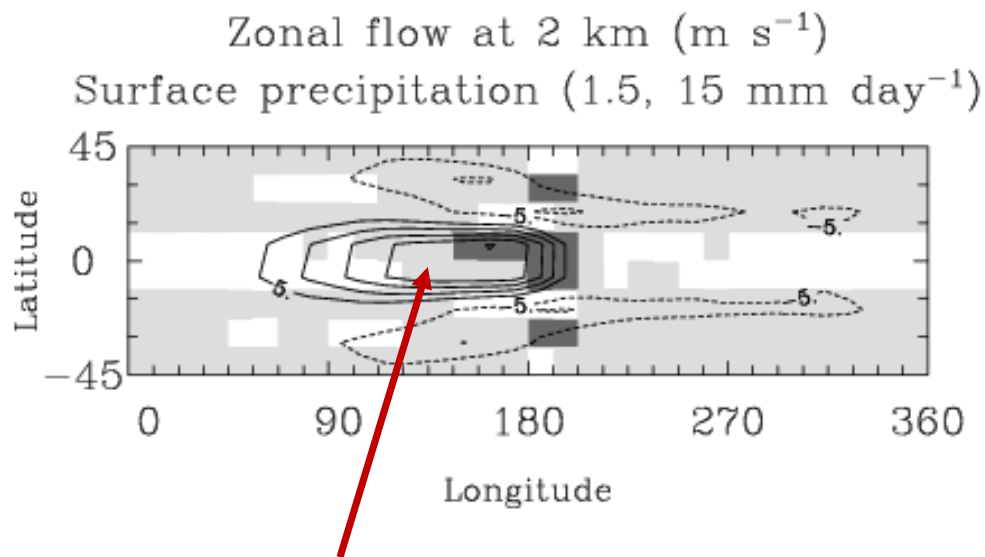
CRCP aligned EW,
free-slip surface,
prescribed radiation

Surface
precipitation

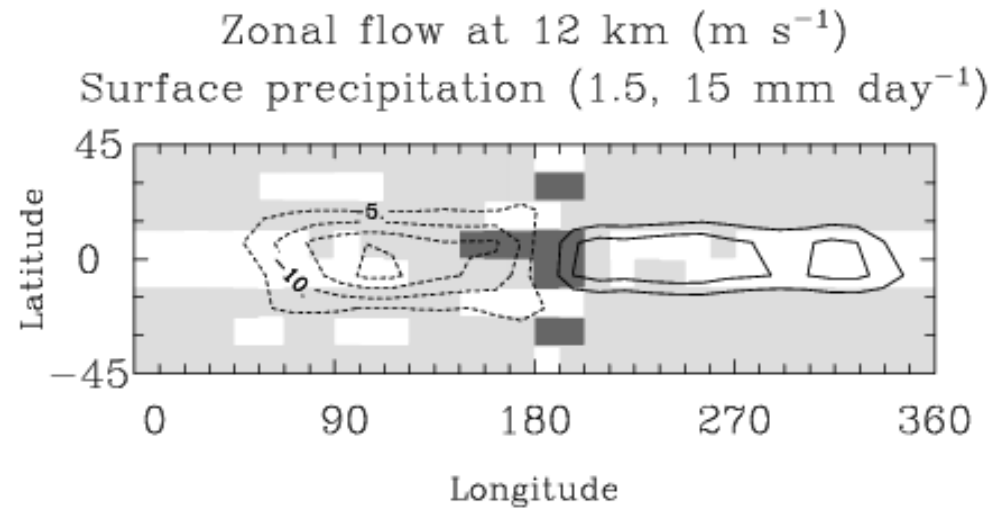


Zonal flow (ground-relative) and surface precipitation,
20-day average in the reference frame moving with
MJO-like coherence

Kelvin/Rossby wave response to east/west

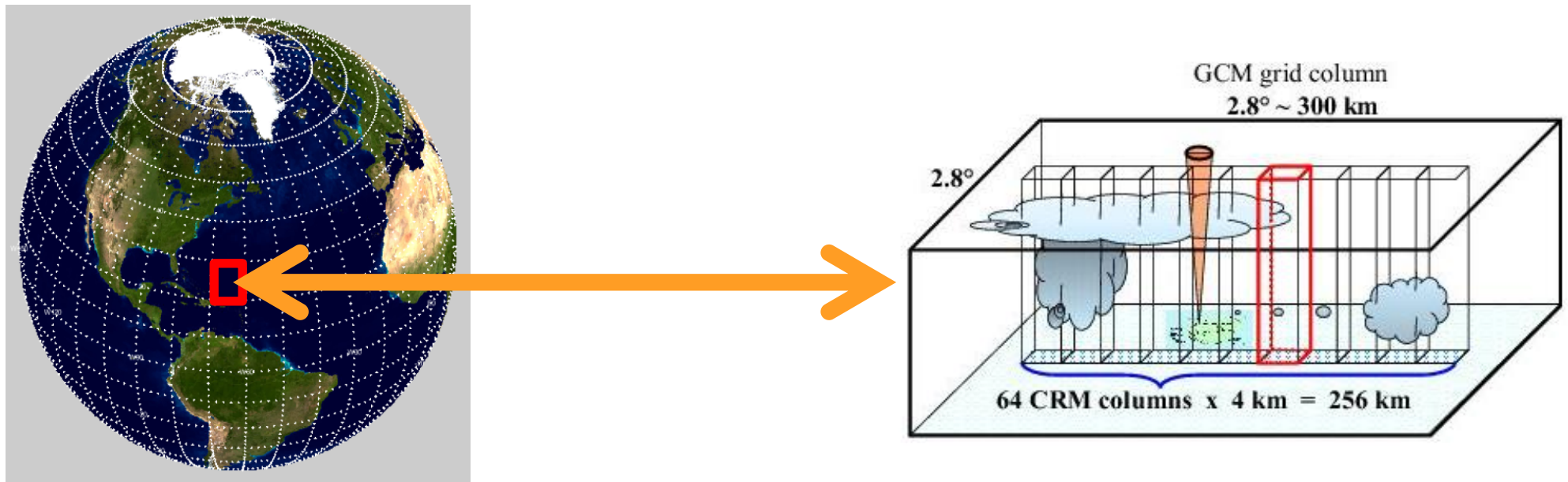


“westerly wind burst”



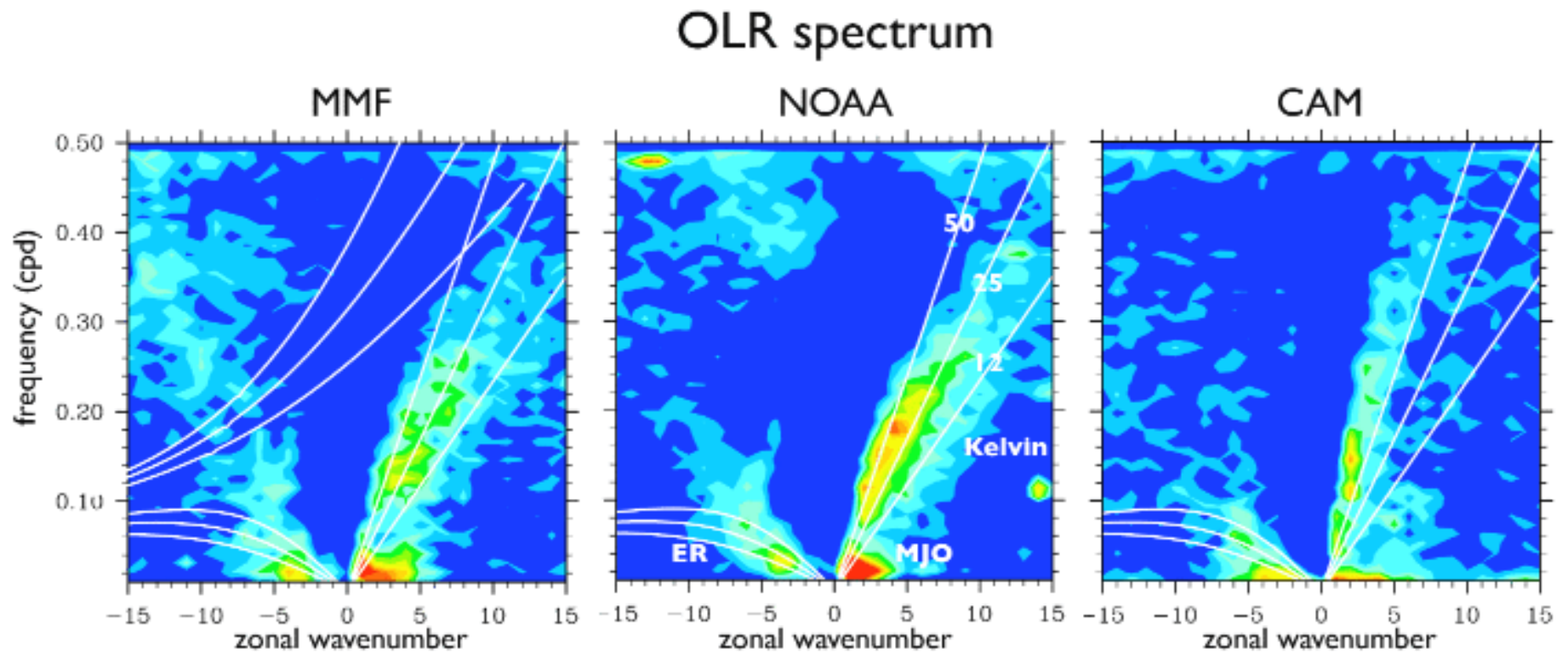
Multiscale Modeling Framework (MMF): SP (Super-Parameterized) CAM (Community Atmospheric Model, part of NCAR's Community Climate System Model (CCSM))

Super-Parameterization



(Khairoutdinov and Randall, 2001; Khairoutdinov et al. 2005, 2007; Wyant et al. 2006... and many many more, including coupled atmosphere-ocean simulations and land-surface model moved into SP, see an impressive list of publications at <http://www.cmmmap.org/research/pubs-ref.html>)

Tropical disturbances in MMF and standard CAM compared to observations on the Wheeler-Kiladis diagram



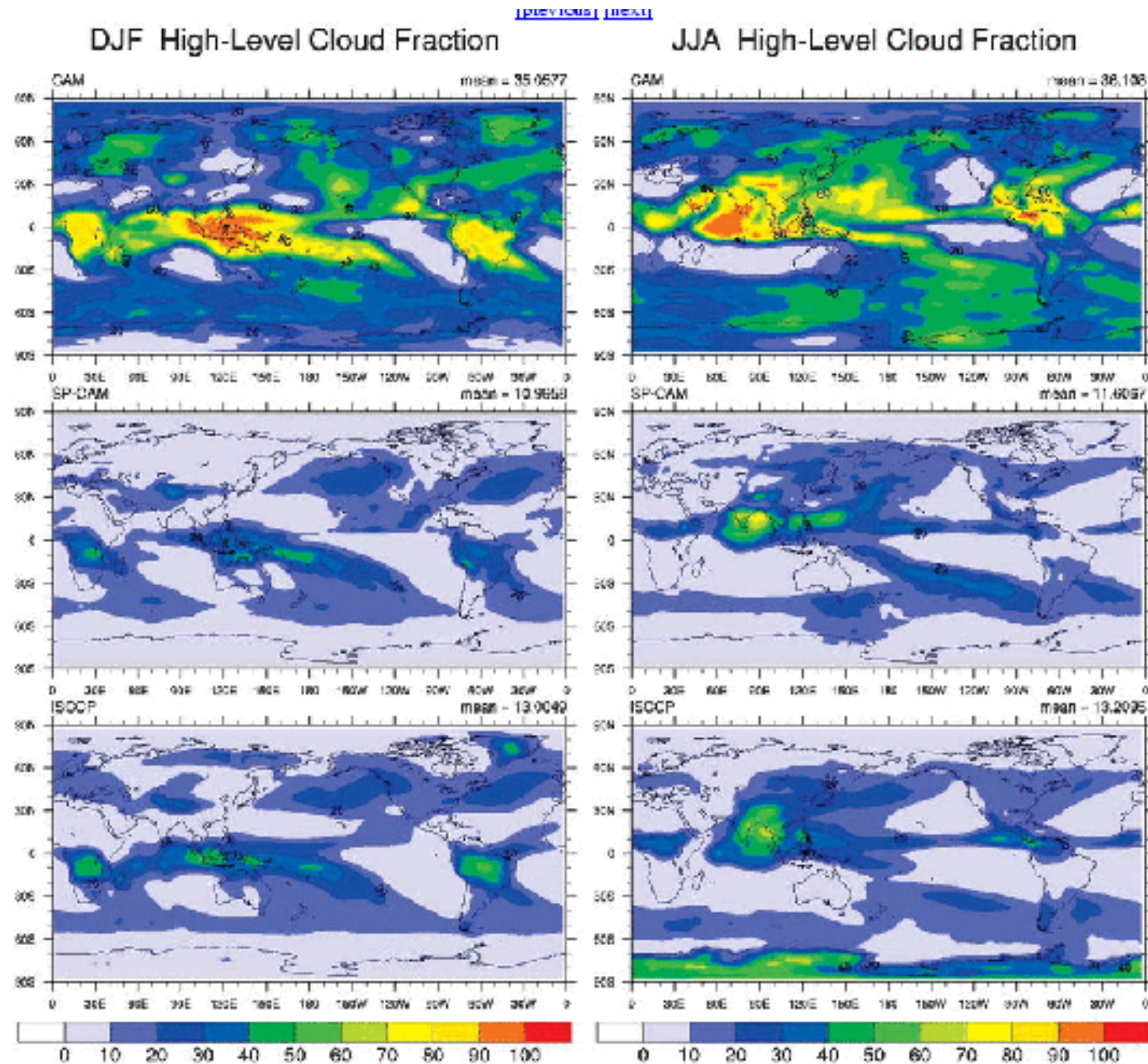
(Khairoutdinov et al. JAS (2007))

Results from a traditional climate model versus MMF

Traditional

MMF

Observations



The works of CMMAP (2006-2016):

- studies of various aspects of intraseasonal variability and MJO;
- including HOC turbulence scheme into embedded CRM;
- development of global CRM;
- expanding atmosphere-only (SP-CAM) simulations to simulations with coupled ocean (ENSO etc.);
- simulations with land-surface model embedded within CRM;
- development of a next-generation of SP model.

about 400 peer-reviewed publications

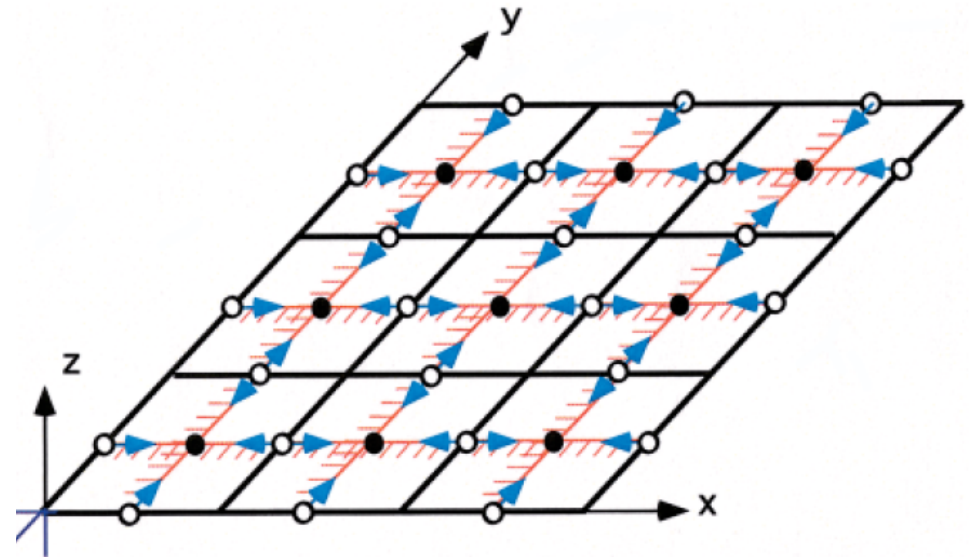
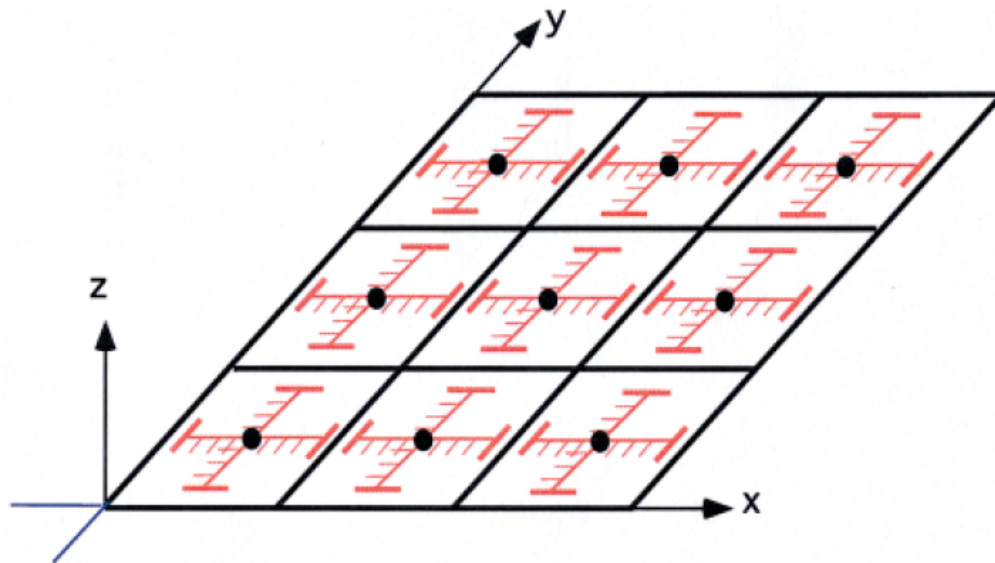
<http://saddleback.atmos.colostate.edu/cmmap/research/pubs-ref.html>

a brief review (and more!) in Grabowski (JMSJ 2016)

BREAKING THE CLOUD PARAMETERIZATION DEADLOCK

BAMS 2003

BY DAVID RANDALL, MARAT KHAIROUTDINOV, AKIO ARAKAWA, AND WOJCIECH GRABOWSKI

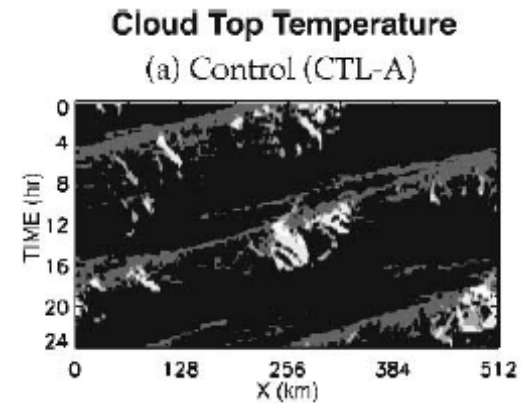
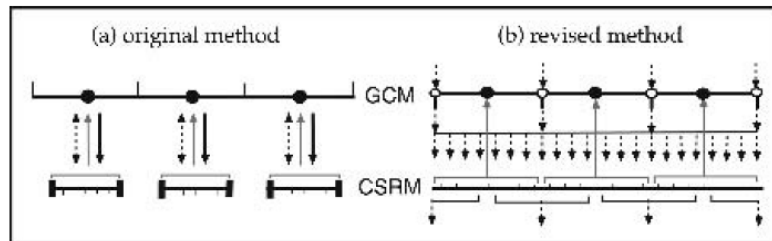


Preliminary Tests of Multiscale Modeling with a Two-Dimensional Framework: Sensitivity to Coupling Methods

MWR 2003

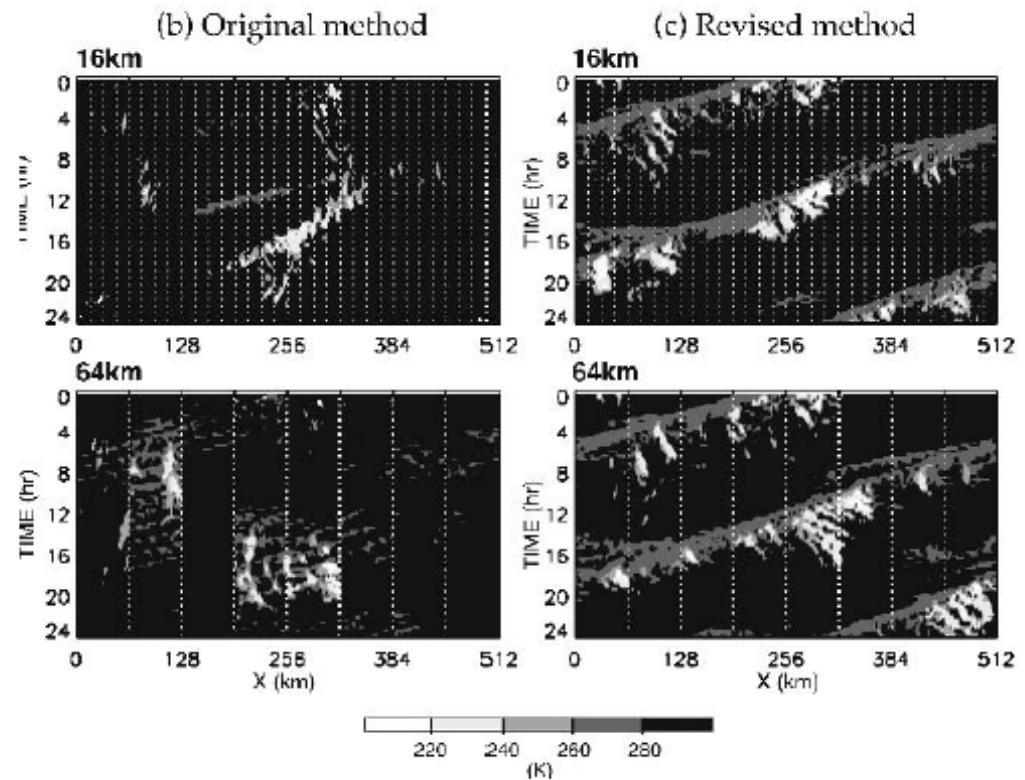
JOON-HEE JUNG* AND AKIO ARAKAWA

Department of Atmospheric Sciences, University of California, Los Angeles, Los Angeles, California

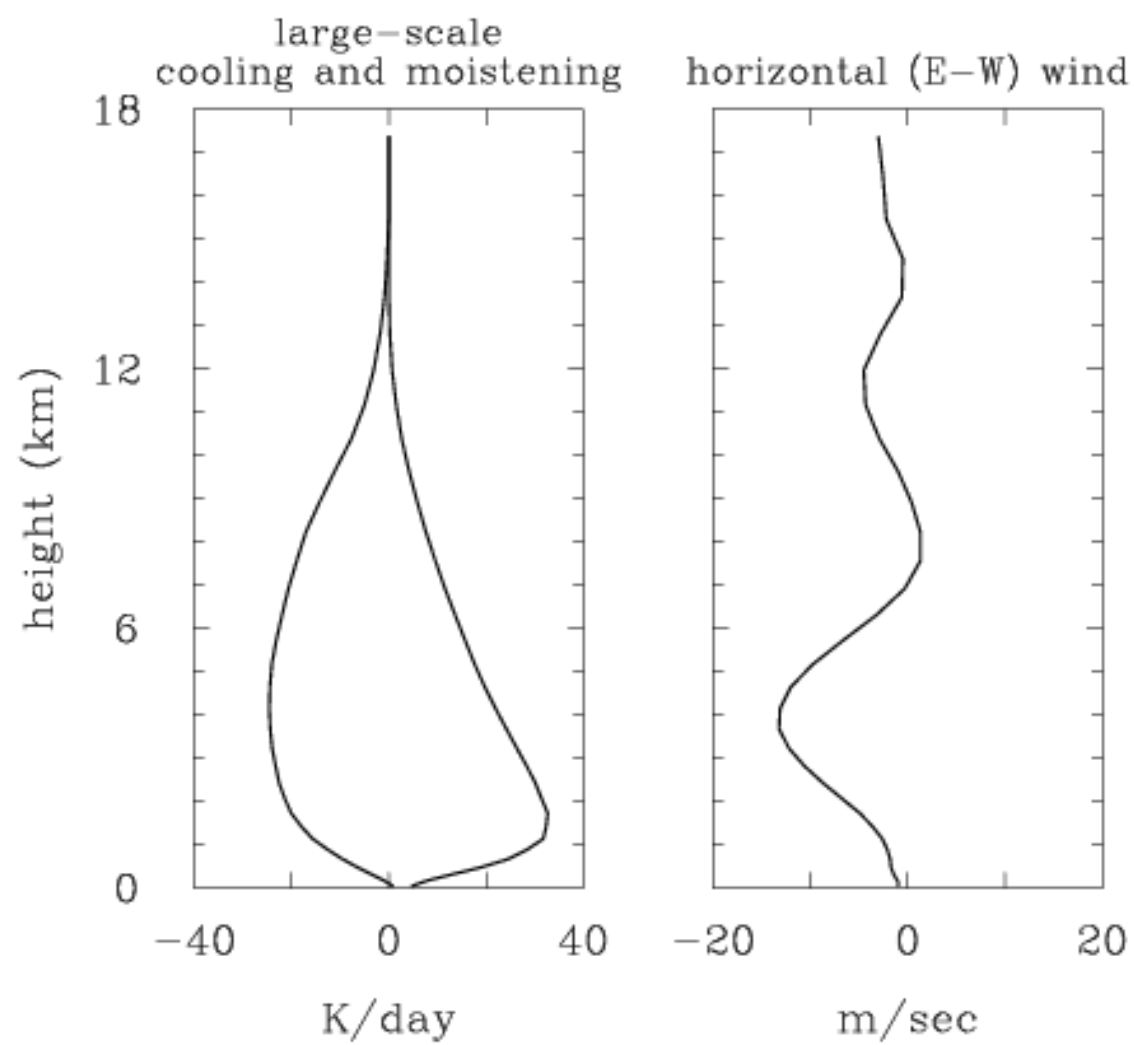


CRM
(benchmark)

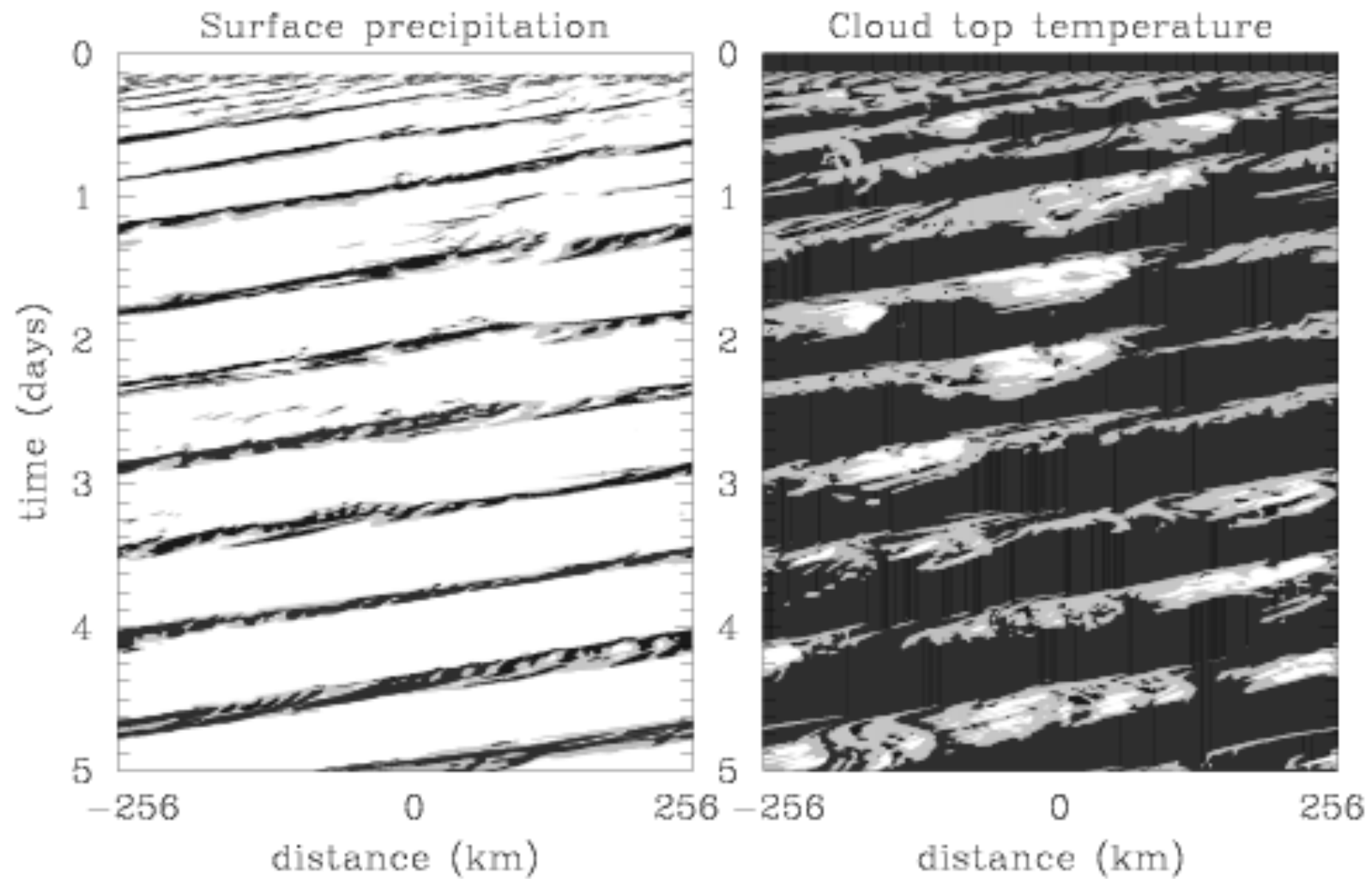
Tests with a strongly
sheared environment



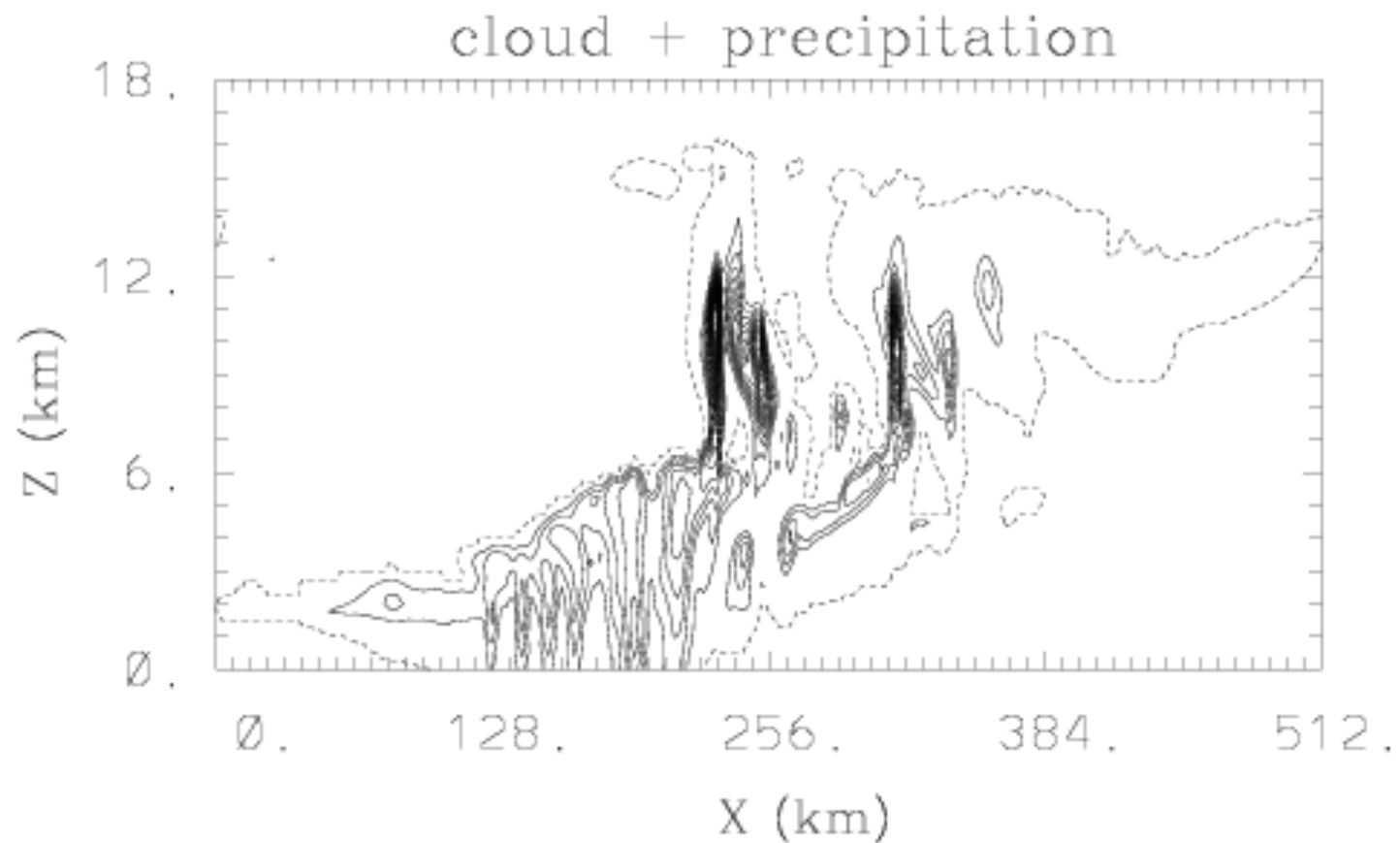
SP



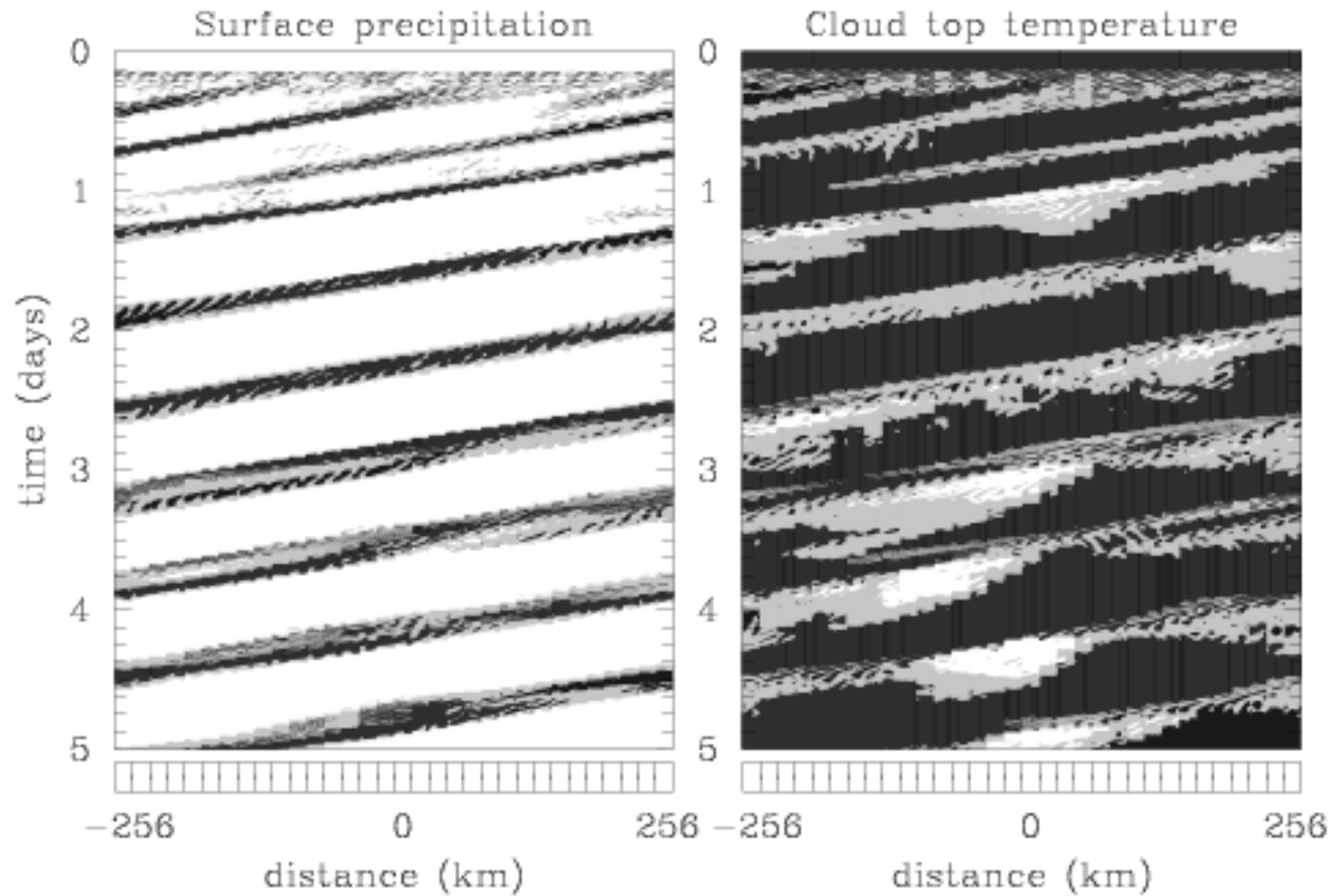
Cloud-resolving simulation (benchmark): $\Delta x=2\text{km}$



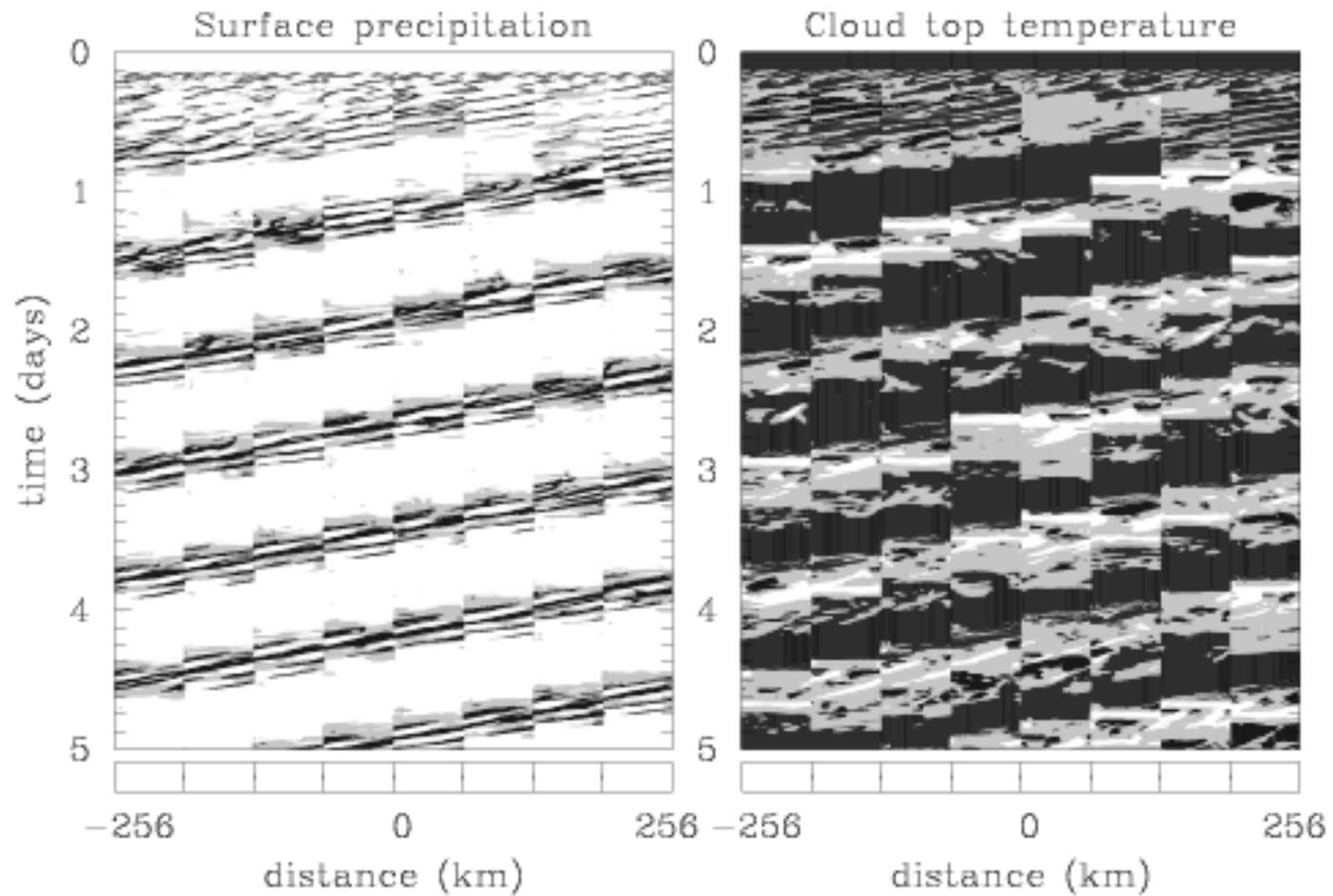
Cloud-resolving simulation (benchmark): $\Delta x=2\text{km}$



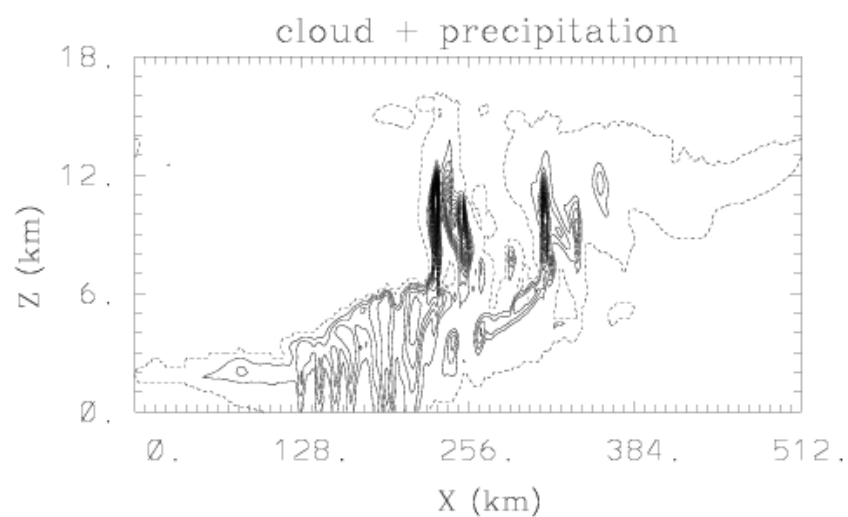
SP simulation: 32 columns with 16-km periodic small-scale models



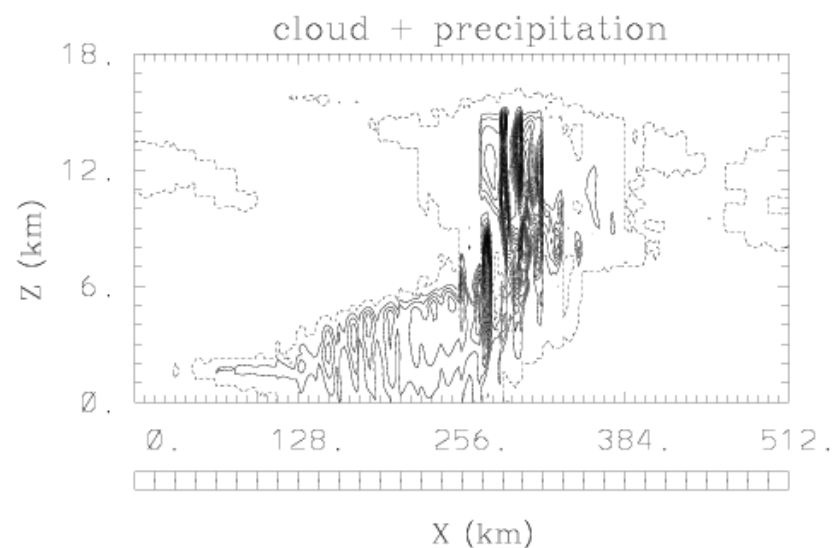
SP simulation: 8 columns with 64-km periodic small-scale models



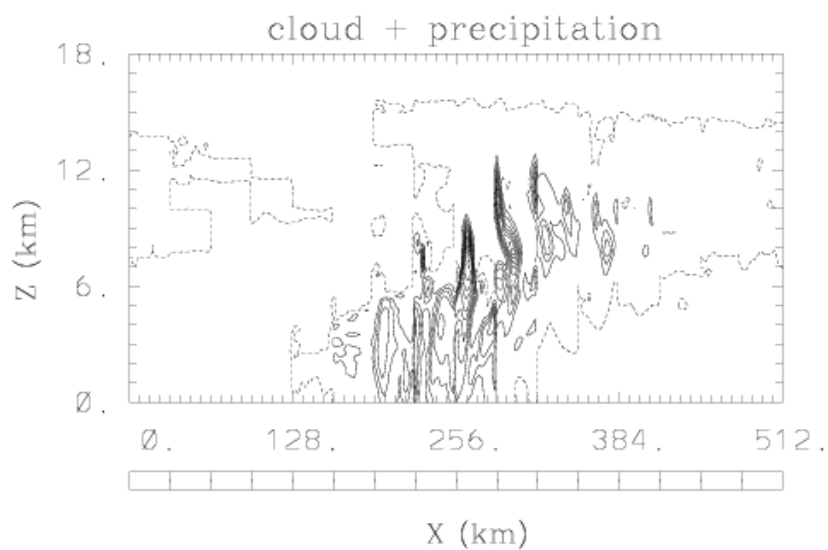
Cloud-resolving simulation (benchmark): $\Delta x=2\text{km}$



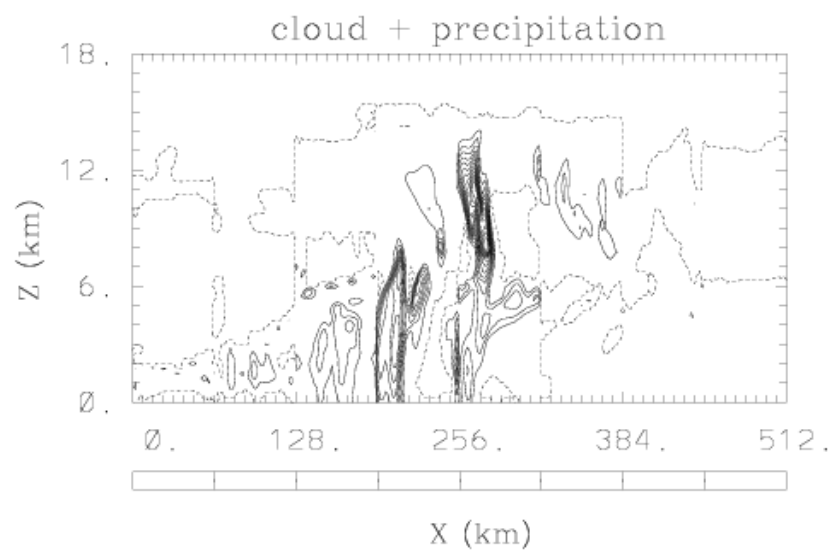
32 columns with 16-km periodic small-scale models



16 columns with 32-km periodic small-scale models

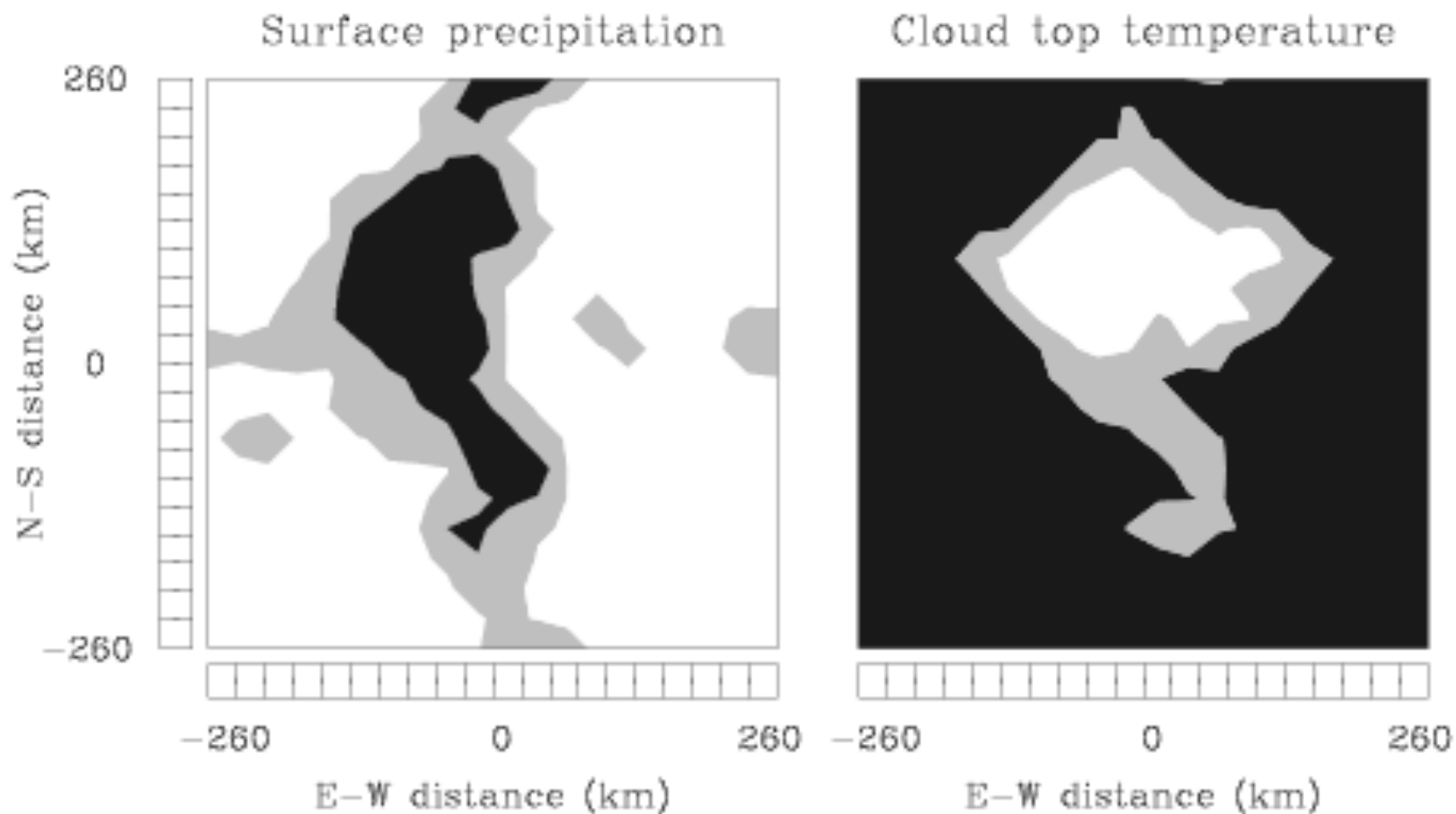


8 columns with 64-km periodic small-scale models

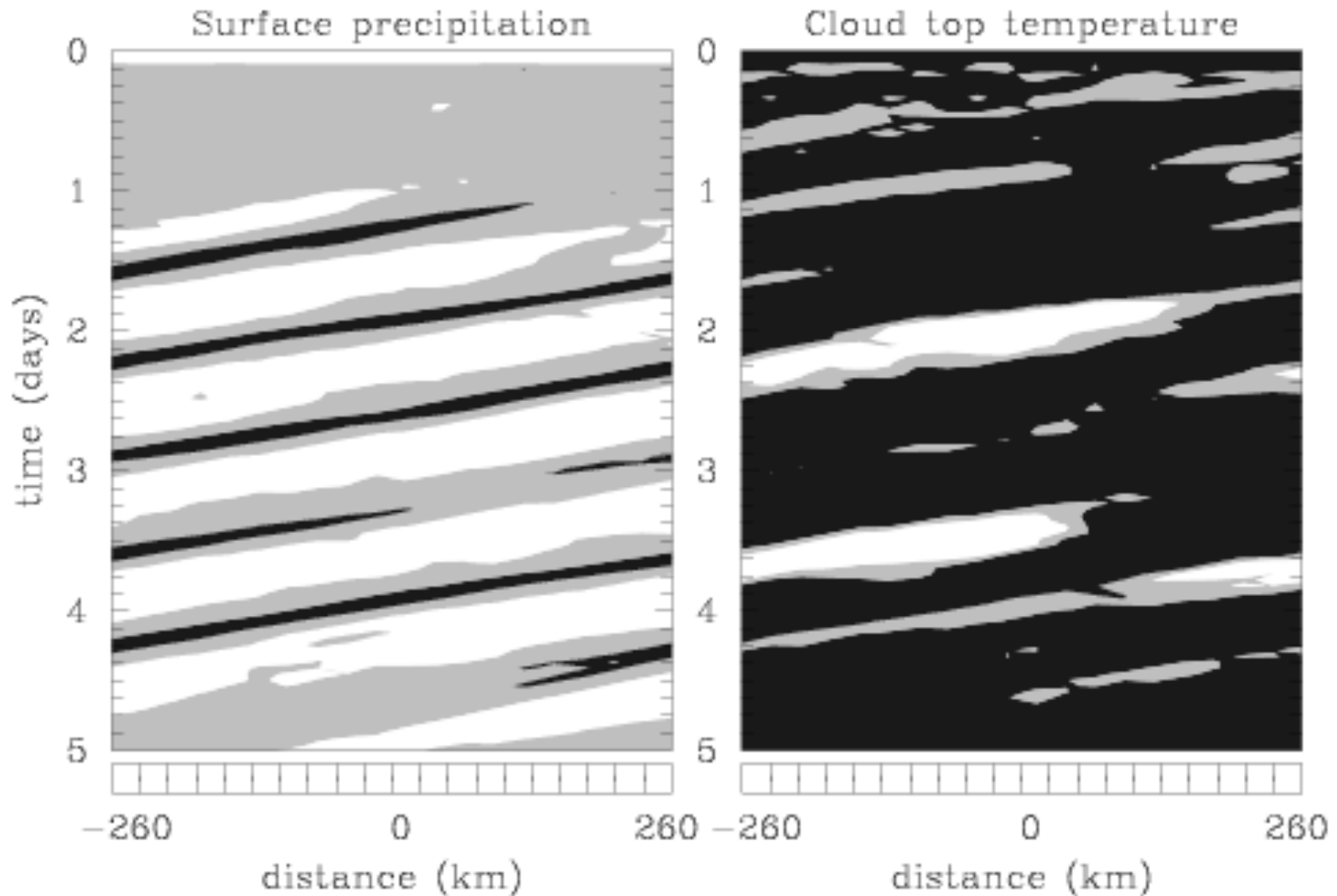


This approach extends naturally into 3D mesoscale model:
2D convective dynamics plus 3D mesoscale dynamics

Snapshots from a 3D simulation in the same setup as before, 520-km mesoscale domain, 26-km grid; 26-km SP domains aligned E-W



Hovmoeller diagrams of N-S averaged surface precipitation and cloud-top temperature from the 3D simulation



My take on these results: Super-parameterization (SP) seems a better-posed approach for limited-area mesoscale models, such as regional climate models, than for temporary general circulation models.

This is because SP in a mesoscale model has to treat only convective-scale dynamics; mesoscale dynamics is left for the 3D mesoscale model.

The work after CMMAP:

- ultra-parameterization (Prof. Mike Pritchard, UC Irvine);
- SP-IFS (Marat at ECMWF, Reading);
- Indian SP-climate model (Marat at IITM, Pune);
- continuation of SP-CAM use at CSU (e.g., CREMIP project)

Towards global large-eddy simulation: Super-parameterization revisited

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NCAR, Boulder, Colorado, USA



Grabowski, W. W., 2016, Towards global large eddy simulation: super-parameterization revisited. *J. Met. Soc. Japan*, **94**, 327-344.

NICAM line-up

✓GCM

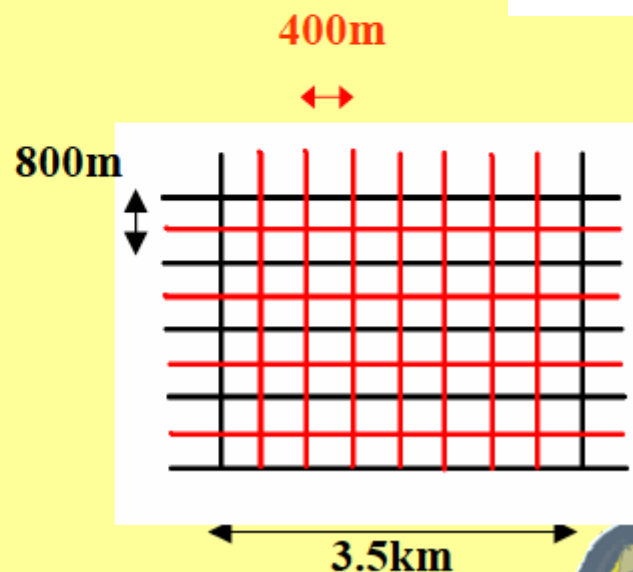
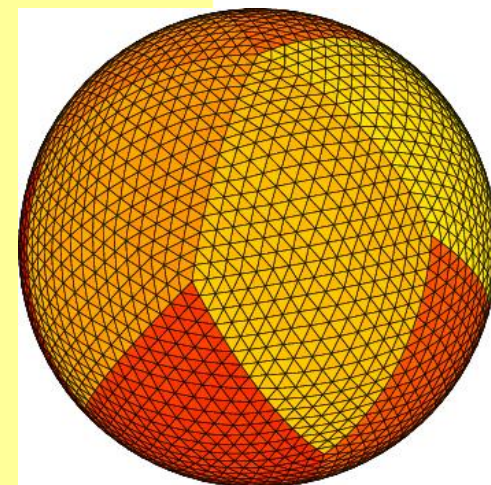
- Glevel 4: 480km
- Glevel 5: 240km
- Glevel 6: 120km
- Glevel 7: 60km
- Glevel 8: 30km

✓GCRM

- Glevel 9: 14km
- Glevel 10: 7km
- Glevel 11: 3.5km

✓GLEM

- Glevel 12: 1.7km
- Glevel 13: 800m
- Glevel 14: 400m

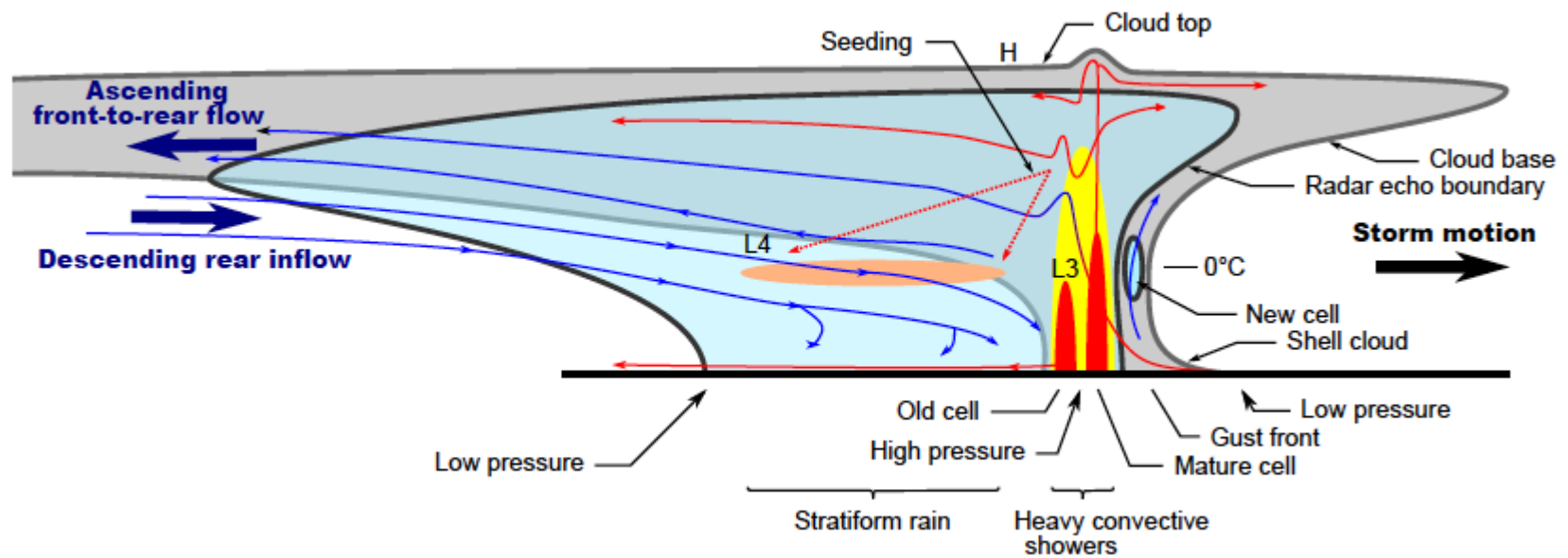
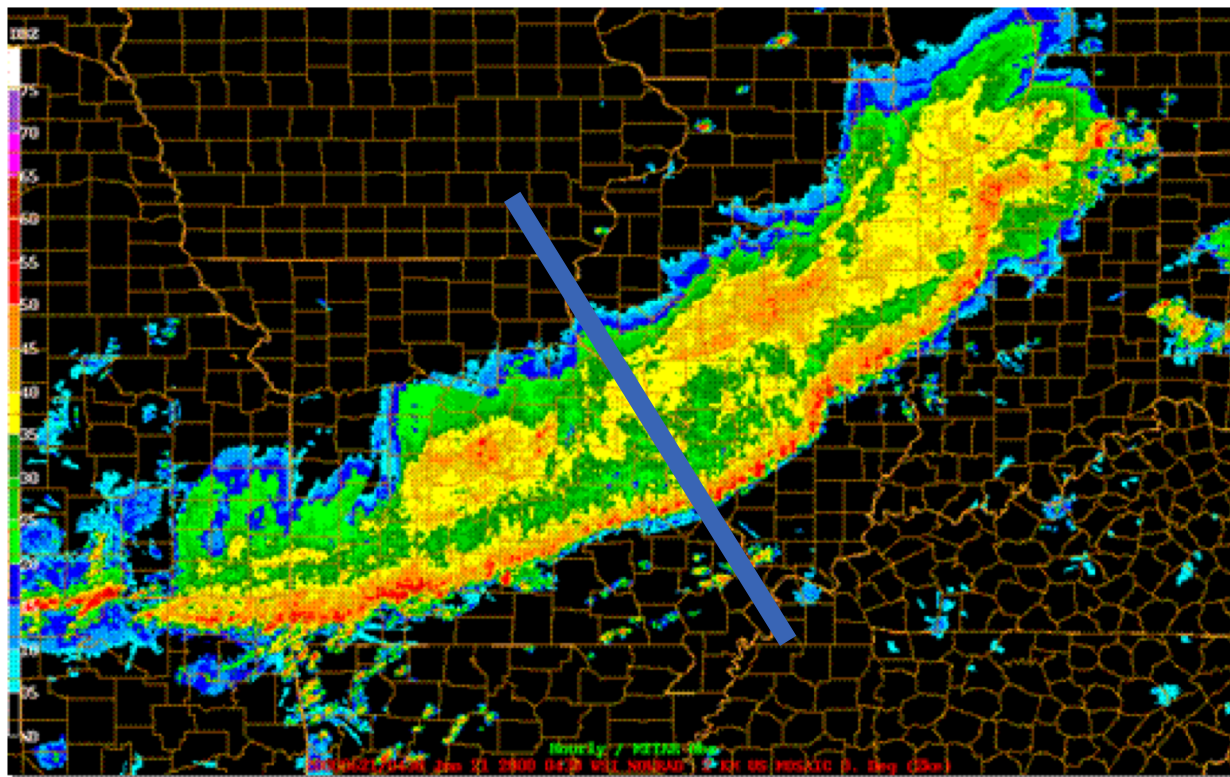


Next Generation Climate Model



Why LES?

Resolution requirements for deep convection...



Resolution Requirements for the Simulation of Deep Moist Convection

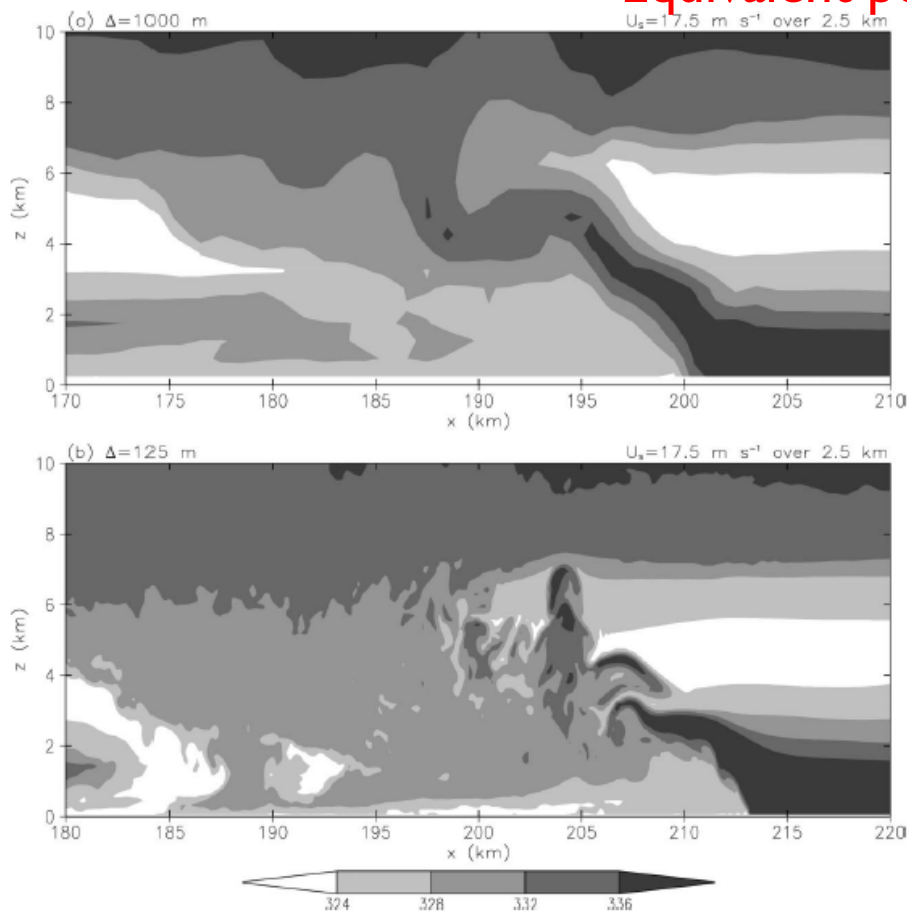
GEORGE H. BRYAN, JOHN C. WYNGAARD, AND J. MICHAEL FRITSCH

Department of Meteorology, The Pennsylvania State University, University Park, Pennsylvania

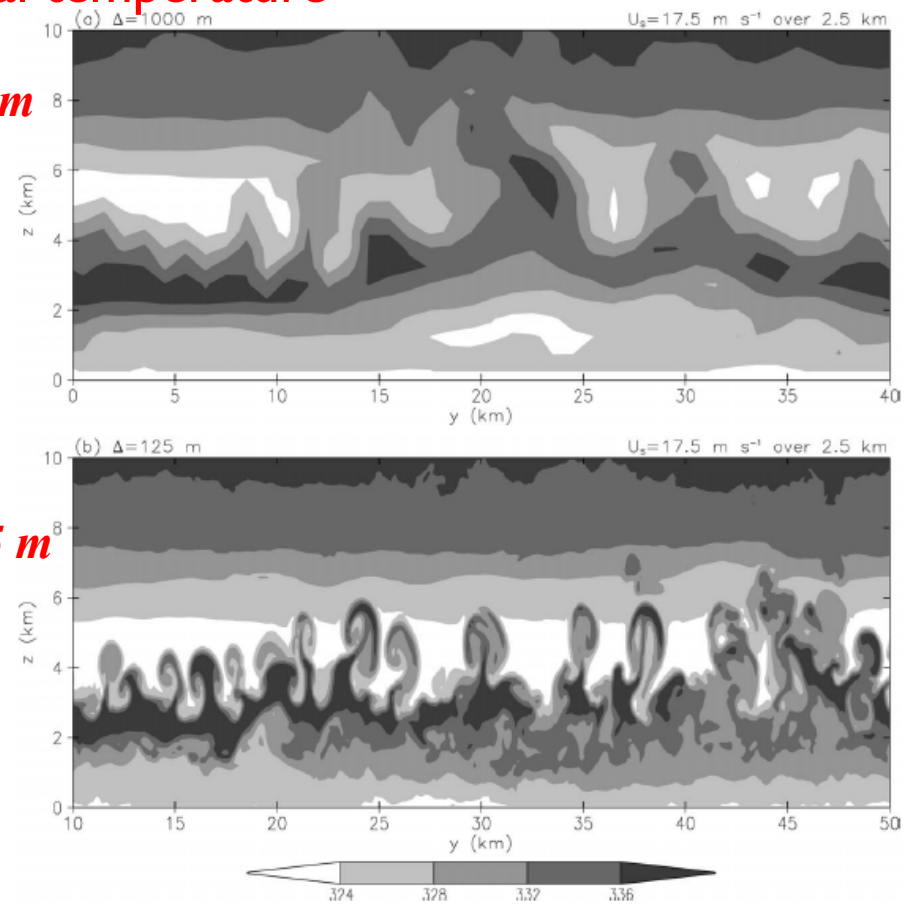
MWR 2003

Squall line simulation:

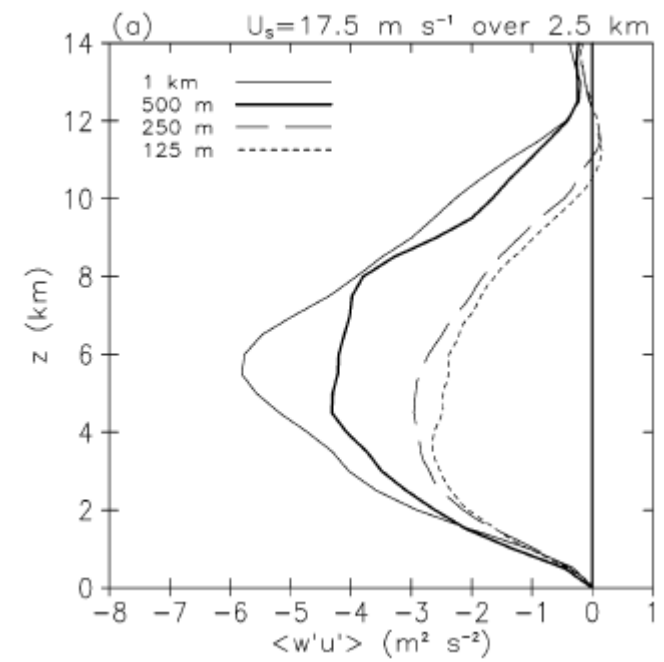
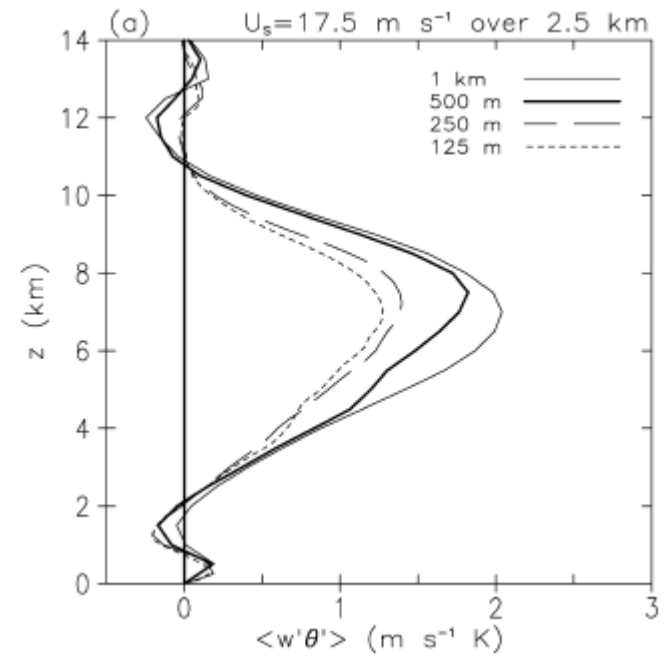
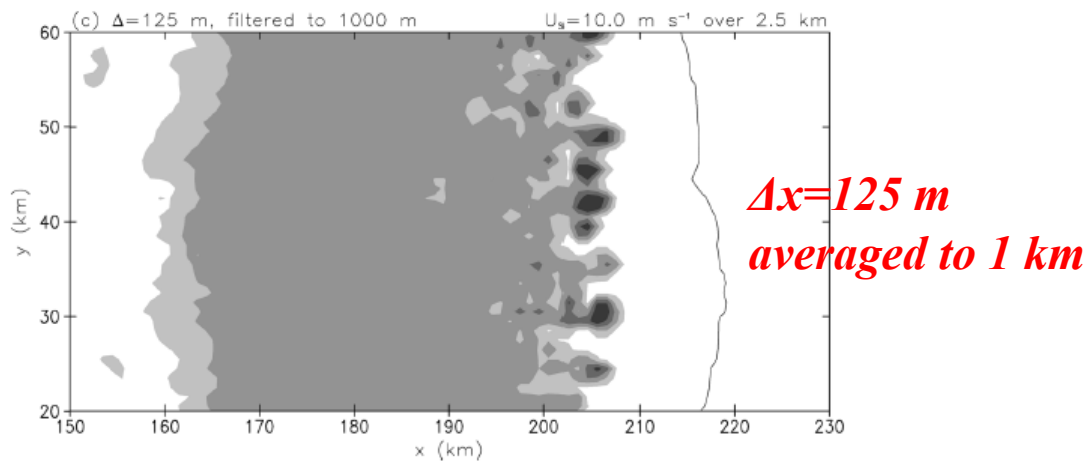
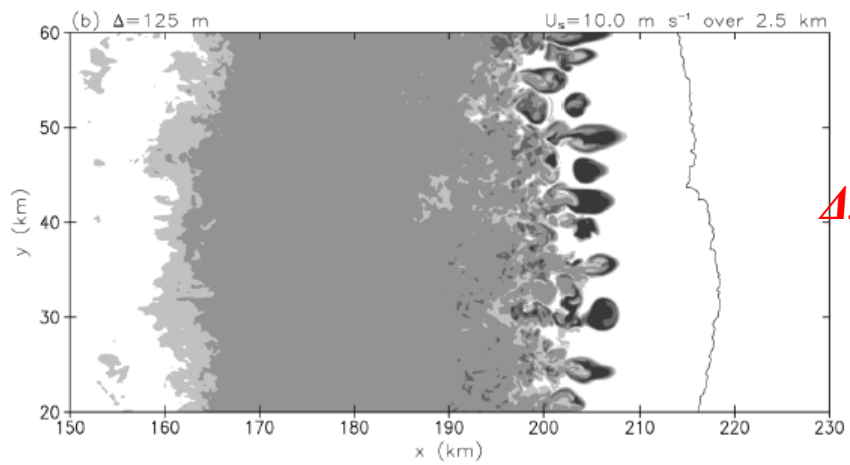
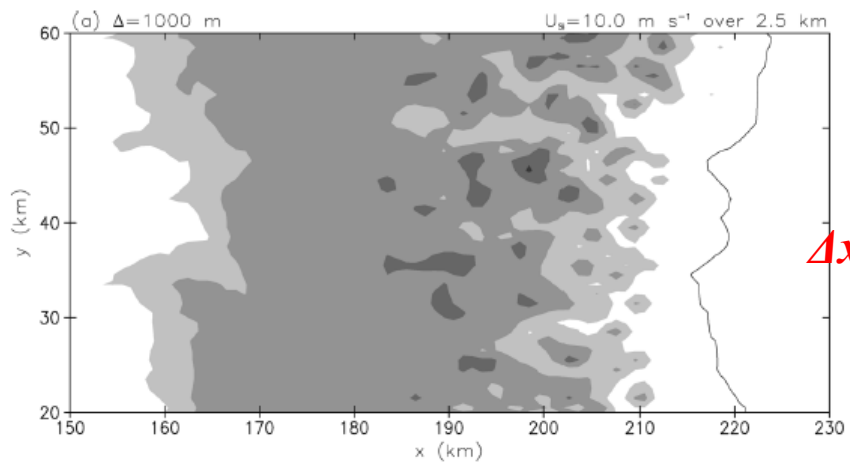
Equivalent potential temperature



Perpendicular to the leading edge



Parallel to the leading edge



Large-Eddy Simulation of Maritime Deep Tropical Convection

Marat F. Khairoutdinov¹, Steve K. Krueger², Chin-Hoh Moeng³, Peter A. Bogenschutz² and David A. Randall⁴



Giga LES

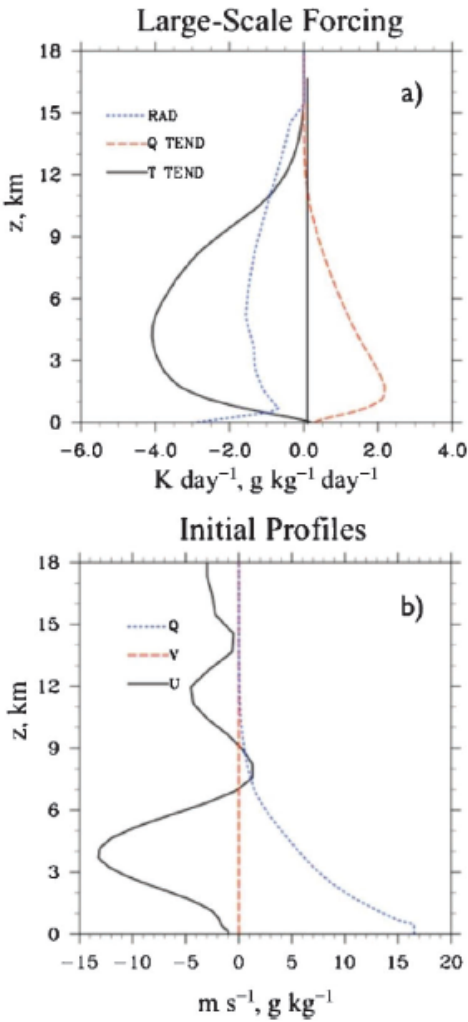
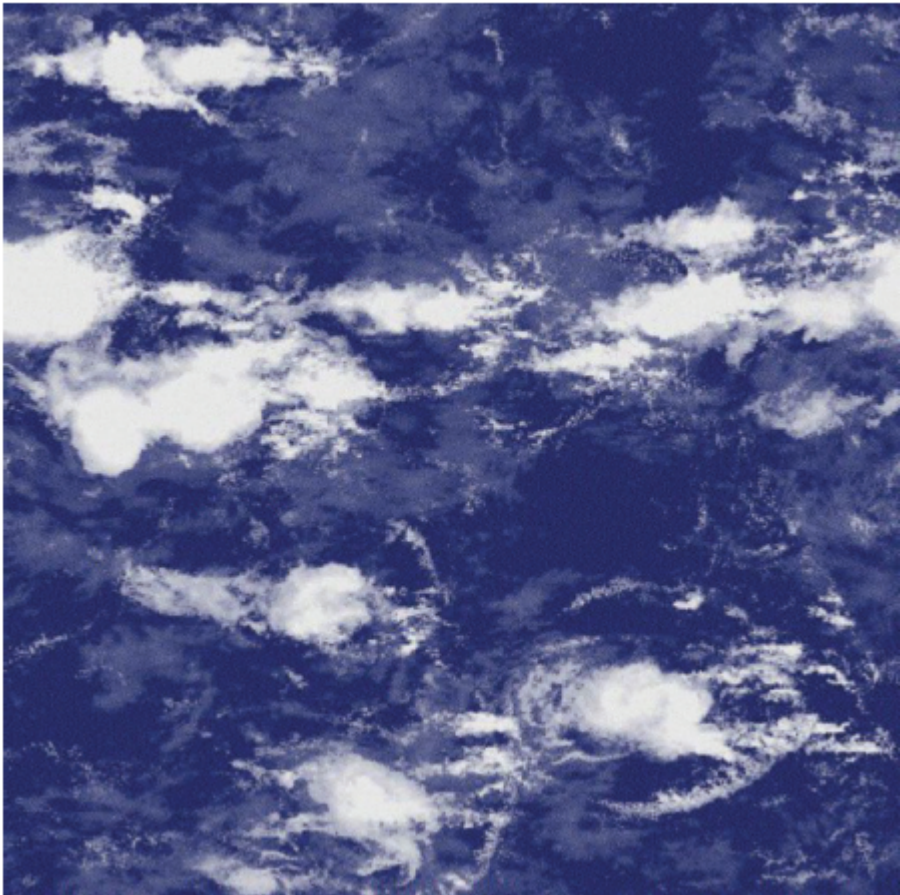


Table 1 Summary of the numerical experiments used in this study.

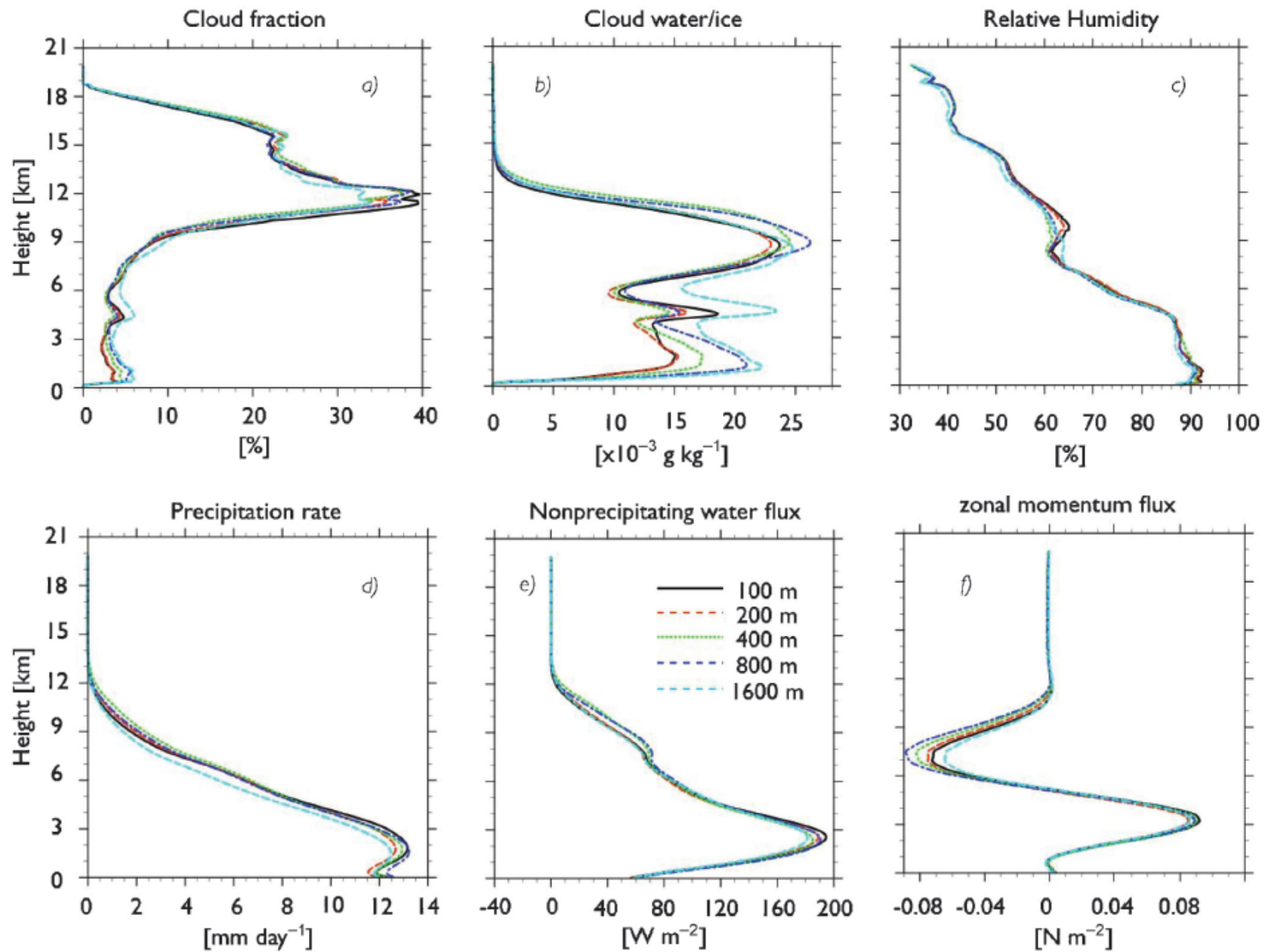
Simulation	Grid size $N_x \times N_y \times N_z$	Horizontal Grid spacing $\Delta x = \Delta y$ (m)	Vertical grid spacing Δz_{\min} - Δz_{\max} (m)
BASE	2048 × 2048 × 256	100	50 - 300
H200	1024 × 1024 × 256	200	50 - 300
H400	512 × 512 × 256	400	50 - 300
H800	256 × 256 × 256	800	50 - 300
H1600	128 × 128 × 256	1600	50 - 300
L64	256 × 256 × 64	800	75 - 500
NOEVP	1024 × 1024 × 256	200	50 - 300



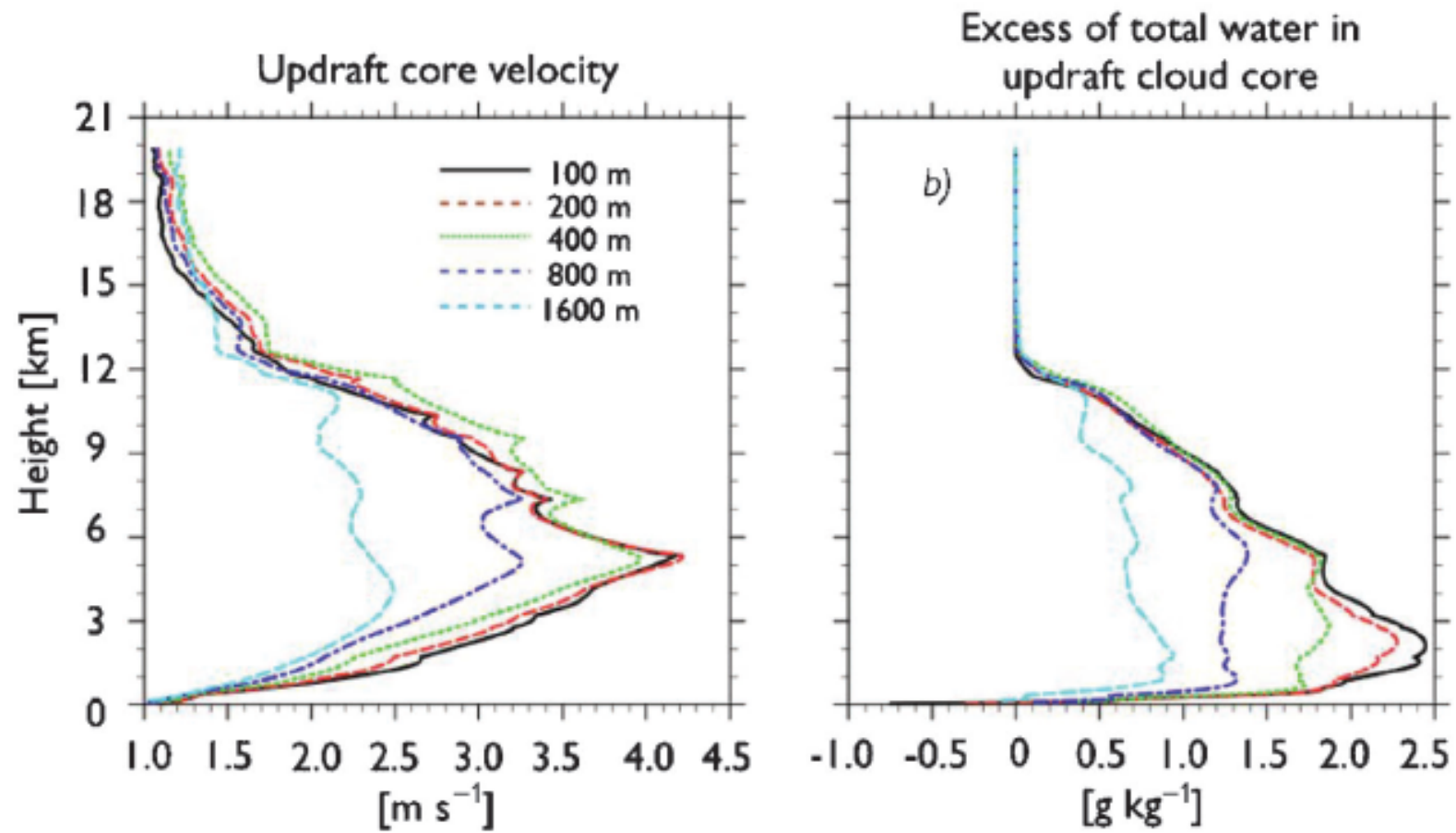
Realistic Giga LES view of deep-convection cloud field



Resolution has a relatively small impact for most bulk fields...

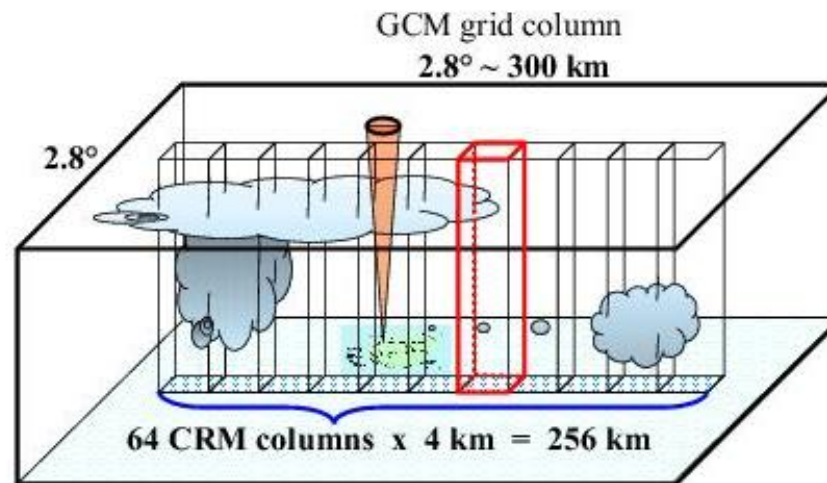


...but the impact is significant for some microphysics-relevant fields:



The original SP applications assumed relatively large outer model domain (100s of km, as in a climate model), implying that both mesoscale and convective dynamics have to be treated in the SP model. **What should be the outer model domain size to capture mesoscale dynamics?**

Think about NWP models in the 80ies...

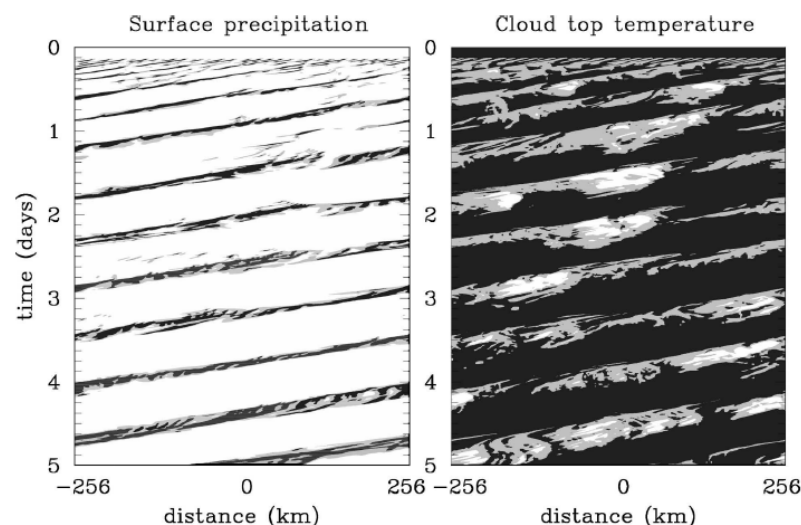


Comments on “Preliminary Tests of Multiscale Modeling with a Two-Dimensional Framework: Sensitivity to Coupling Methods”

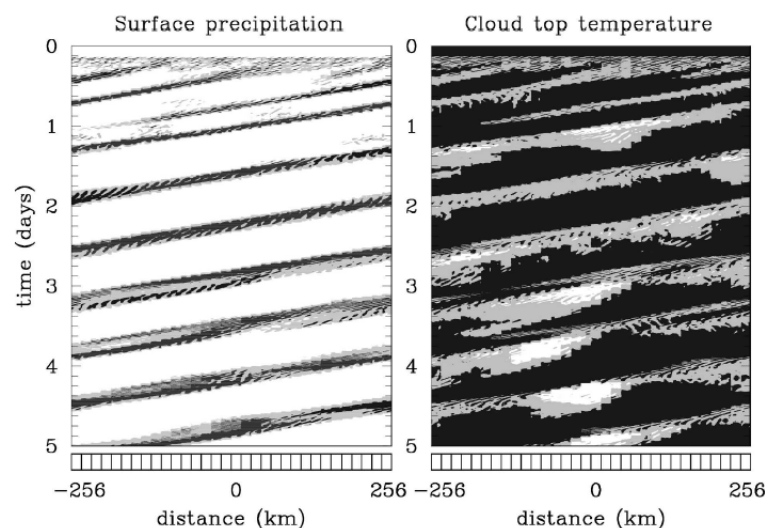
MWR 2006

WOJCIECH W. GRABOWSKI

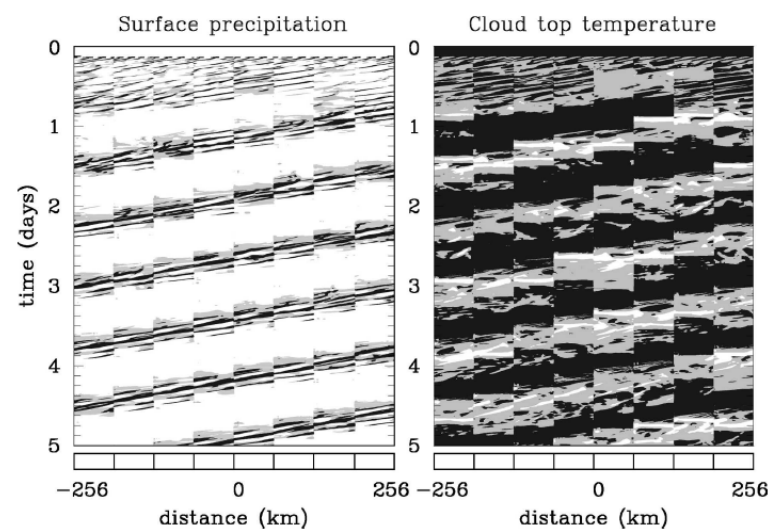
*2D simulations,
 $\Delta x = 2$ km*



CRM
(benchmark)



SP with 16 km domains



SP with 64 km domains

*Natural extension to a
3D outer model:*

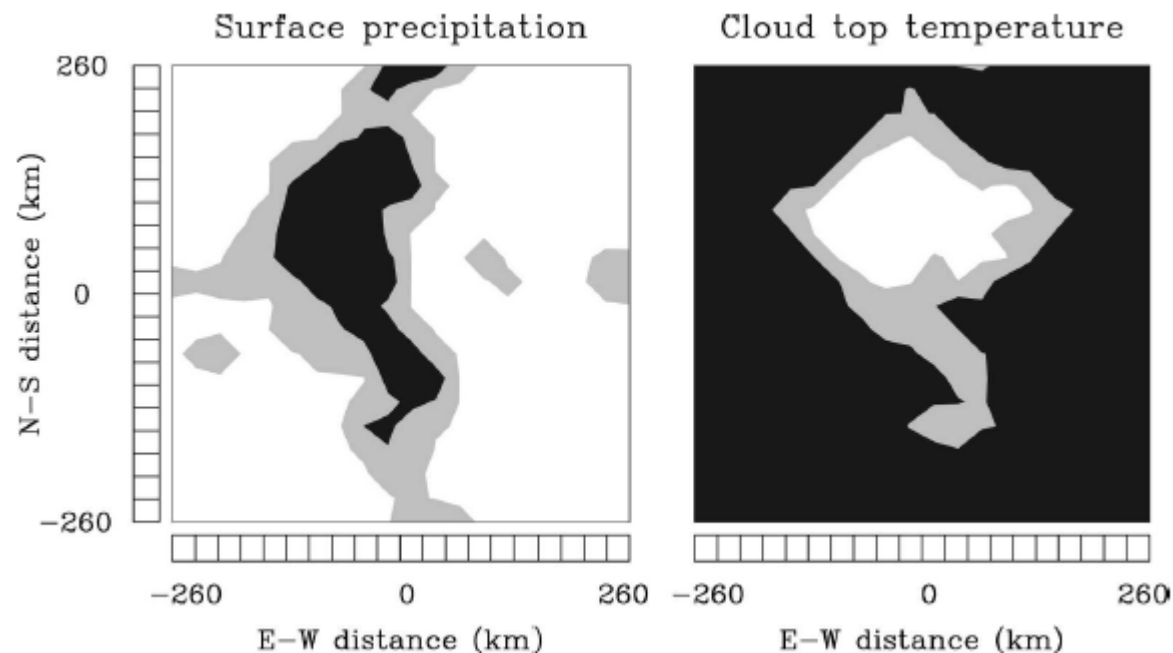
outer model:

$\Delta x = \Delta y = 26 \text{ km}$

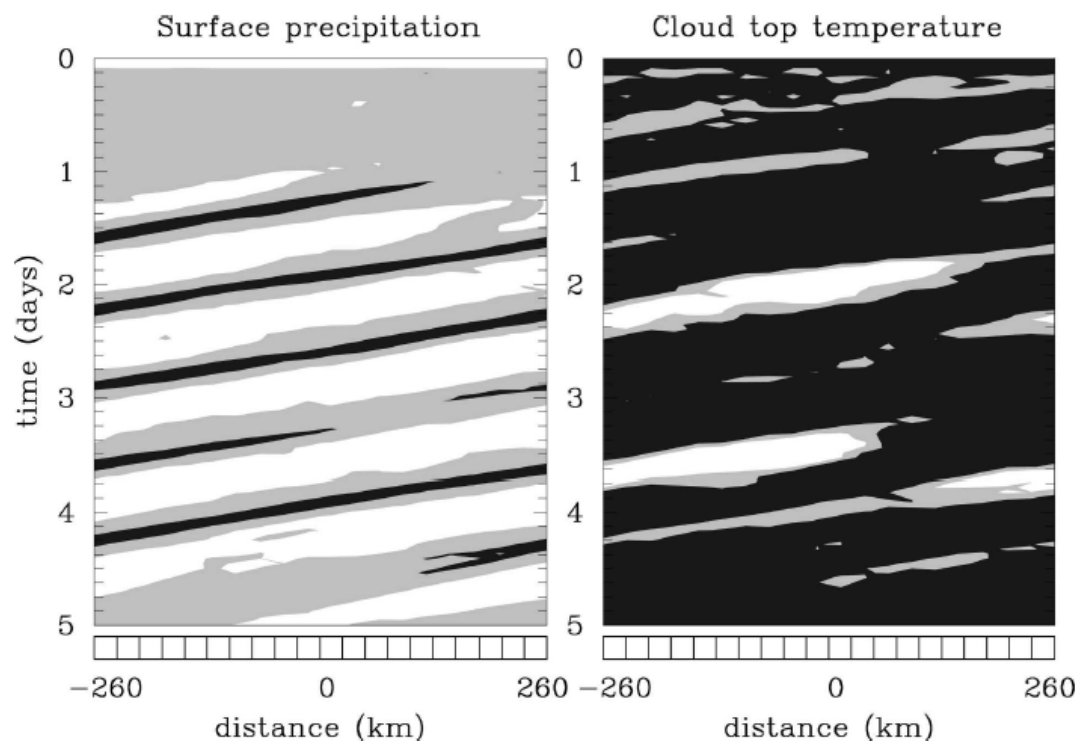
2D SP models

(aligned E-W) with

$\Delta x = 2 \text{ km}$

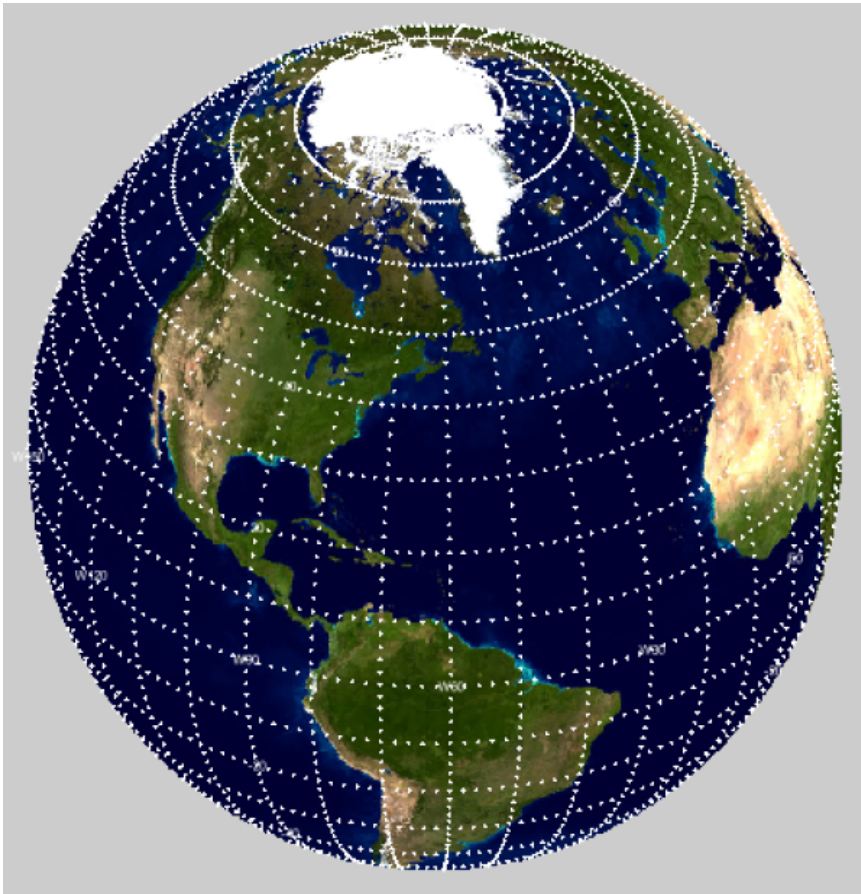


snapshot



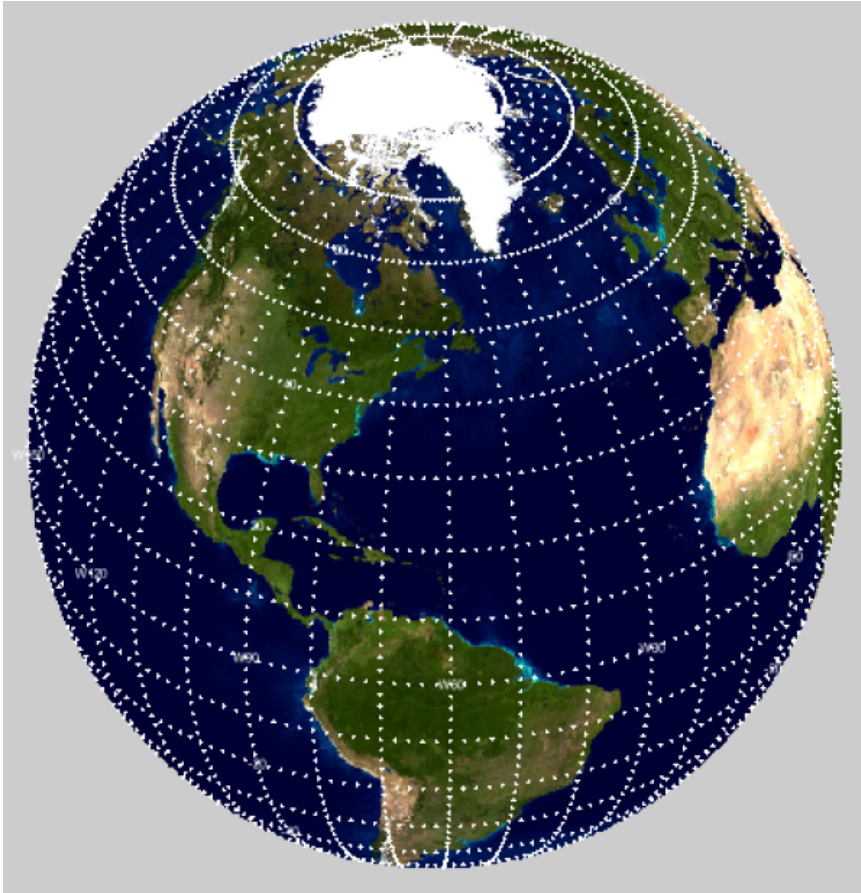
Hovmueller diagram
of N-S averaged fields

If the outer model has a horizontal grid length around a **few tens of km**, it will faithfully represent **mesoscale dynamics**, like 20th century NWP models. The **embedded SP models** need only to cope with small-scale processes, such as **convective-scale dynamics**. They can be **2D** as in the examples above, but they can be **3D**, and **even LES** if boundary layer dynamics or shallow convection is to be well simulated...



Radius: $R \approx 6.4 \times 10^3 \text{ km}$

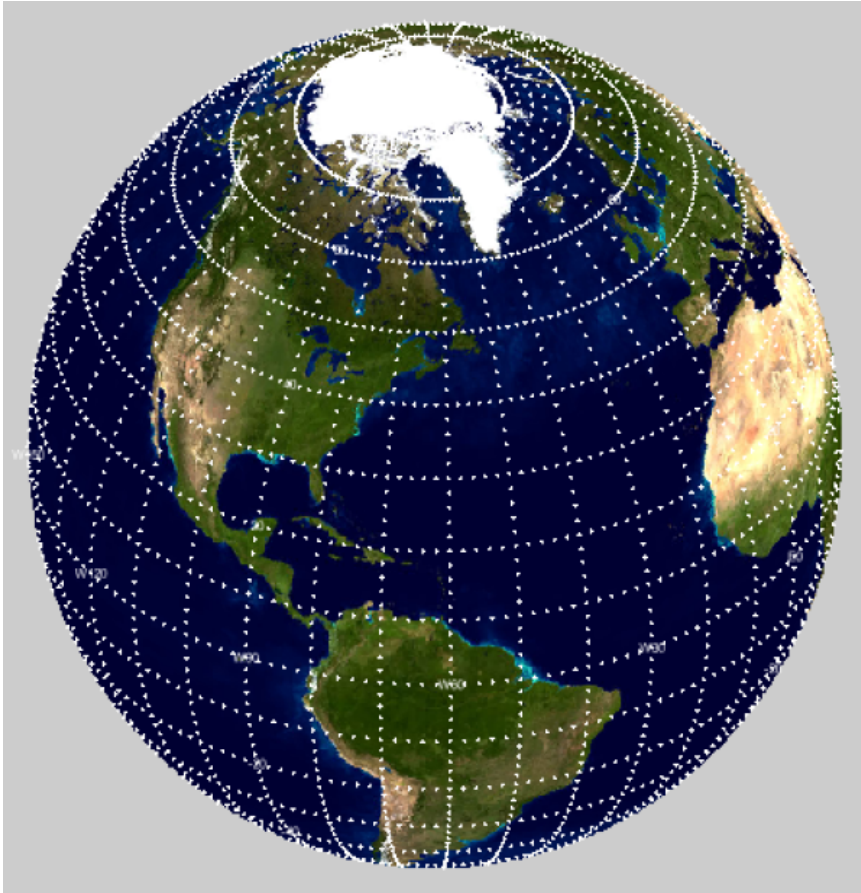
Surface area: $S \approx 5.1 \times 10^8 \text{ km}^2$



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If one would like to cover the surface with **LES squares of 20 km by 20 km**, there will be around **1.3 million squares...**



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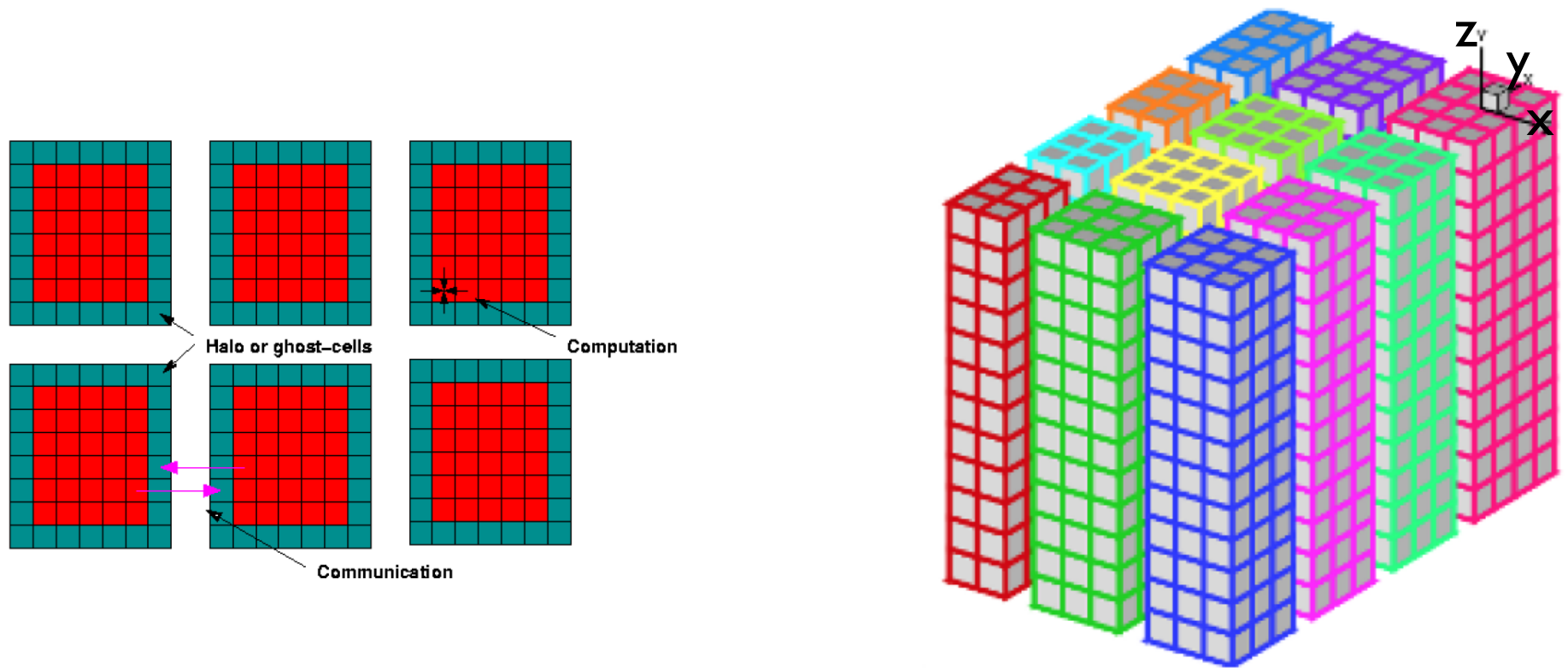
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If one would like to cover the surface with **LES squares** of 20 km by 20 km, there will be around **1.3 million** squares... This suggests that one can apply a computer with up to 1.3 million processors for parallel simulations...

Issues:

- Parallel processing?
- What equations to use?

Domain decomposition for the finite-difference parallel processing



Large amount of data needs to be exchange at every time step in the halos at the sub-domain boundaries. This makes the parallel processing difficult.

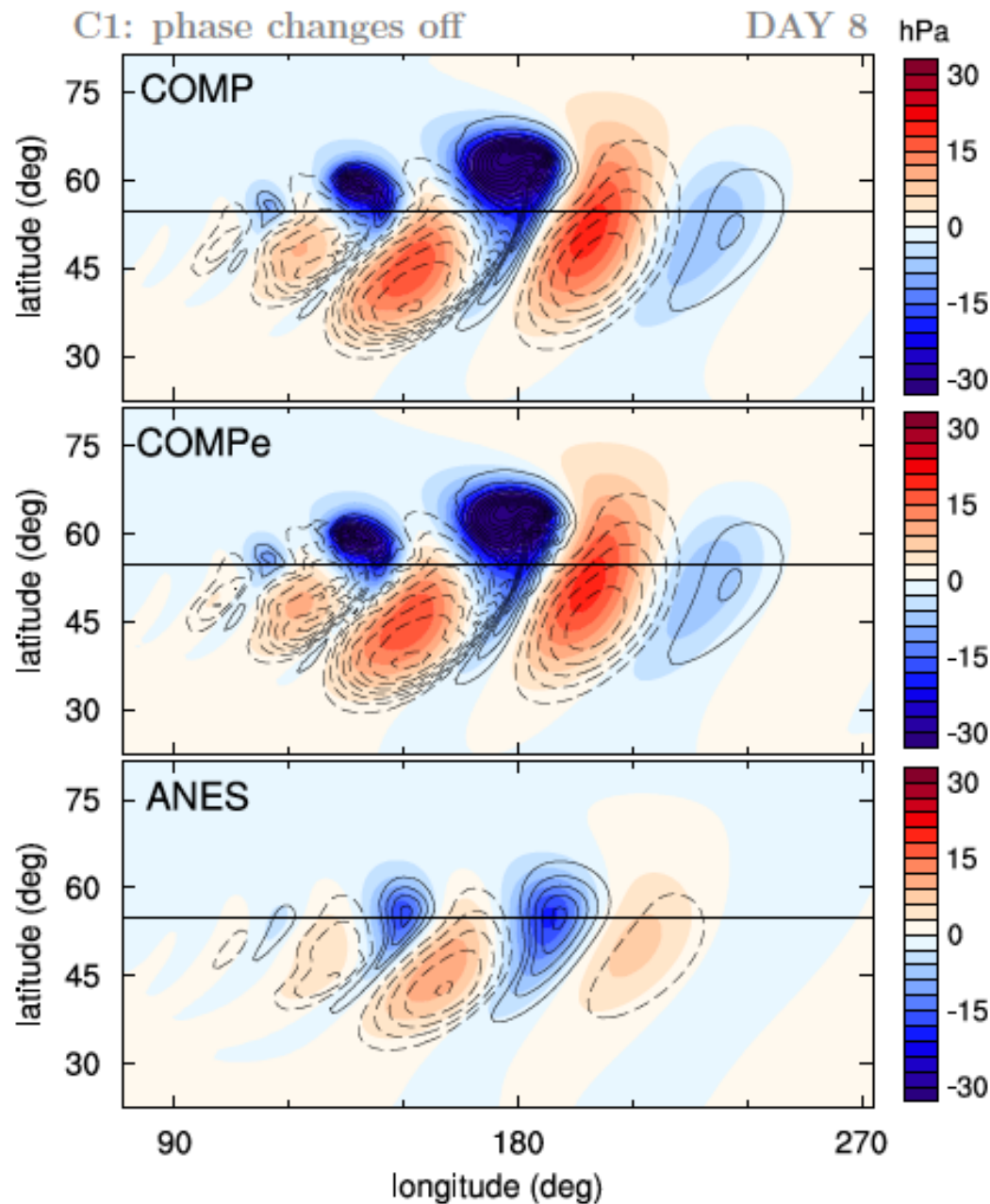
What governing equations to use?

Extension of the small-scale nonhydrostatic equations to the global scale is not trivial.

Compressible dynamics is valid across all scales, but it is numerically cumbersome due to presence of pesky sound waves that can be argued irrelevant for weather and climate.

Anelastic equations are appropriate for small-scale and mesoscale dynamics, but validity of its extension to the global scale is questionable.

Jablonowski and Williamson (2006) baroclinic wave test:



**Implicit compressible
model with $\Delta t = 300$ s**

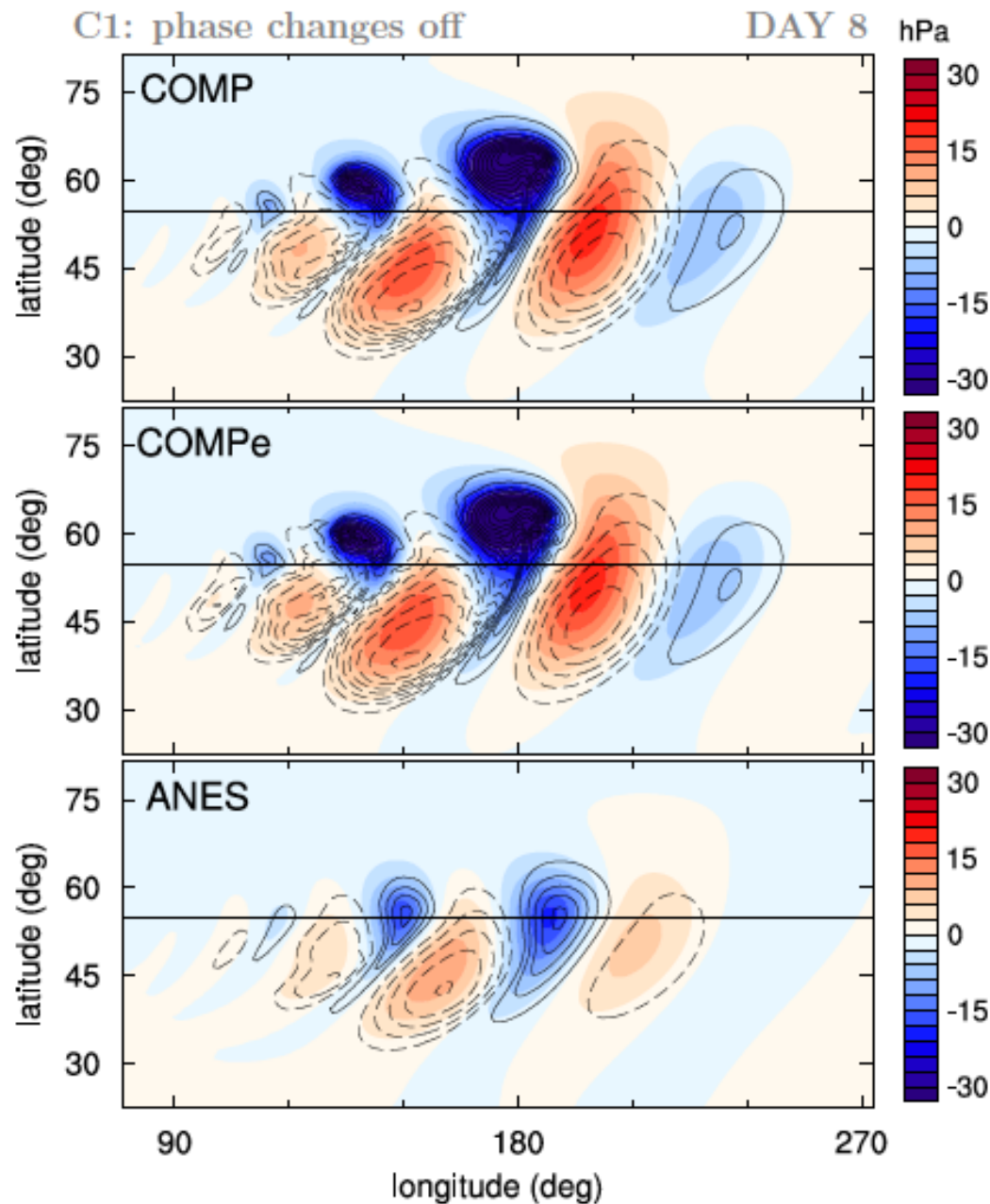
**Explicit
compressible model
with $\Delta t = 2$ s**

**Anelastic model
with $\Delta t = 300$ s**

Surface virtual temperature (contours)
and pressure perturbations (colors).

Smolarkiewicz et al. *JCP* 2014
Kurowski et al. *JAS* 2015

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Conclusions:

- **Anelastic equations are not appropriate for global scales;**
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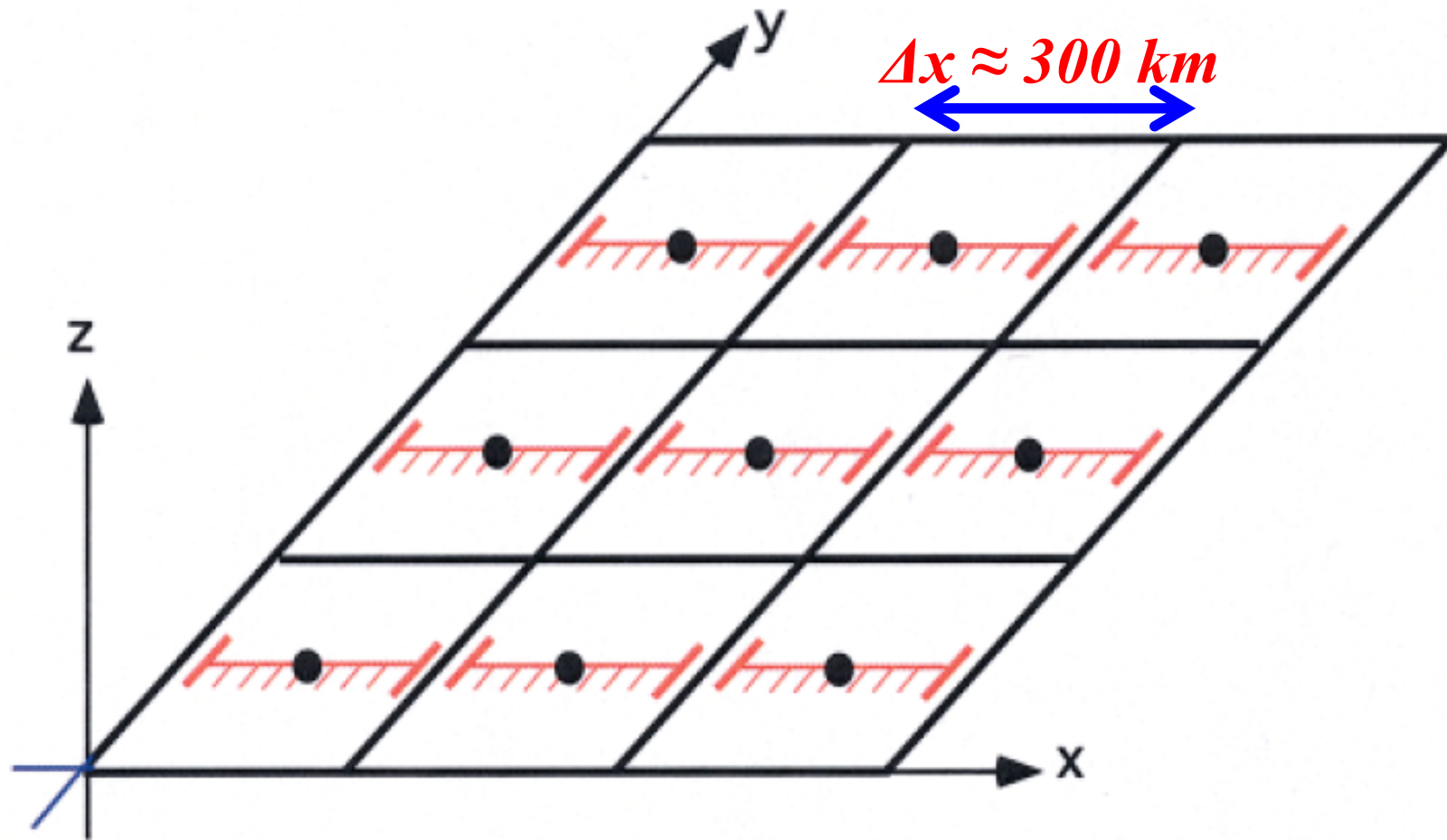
However, pressure solver in the implicit compressible model (significantly more cumbersome than in the anelastic system, see Smolarkiewicz et al. *JCP* 2014) would need to work really hard when global LES is the target...

Issues:

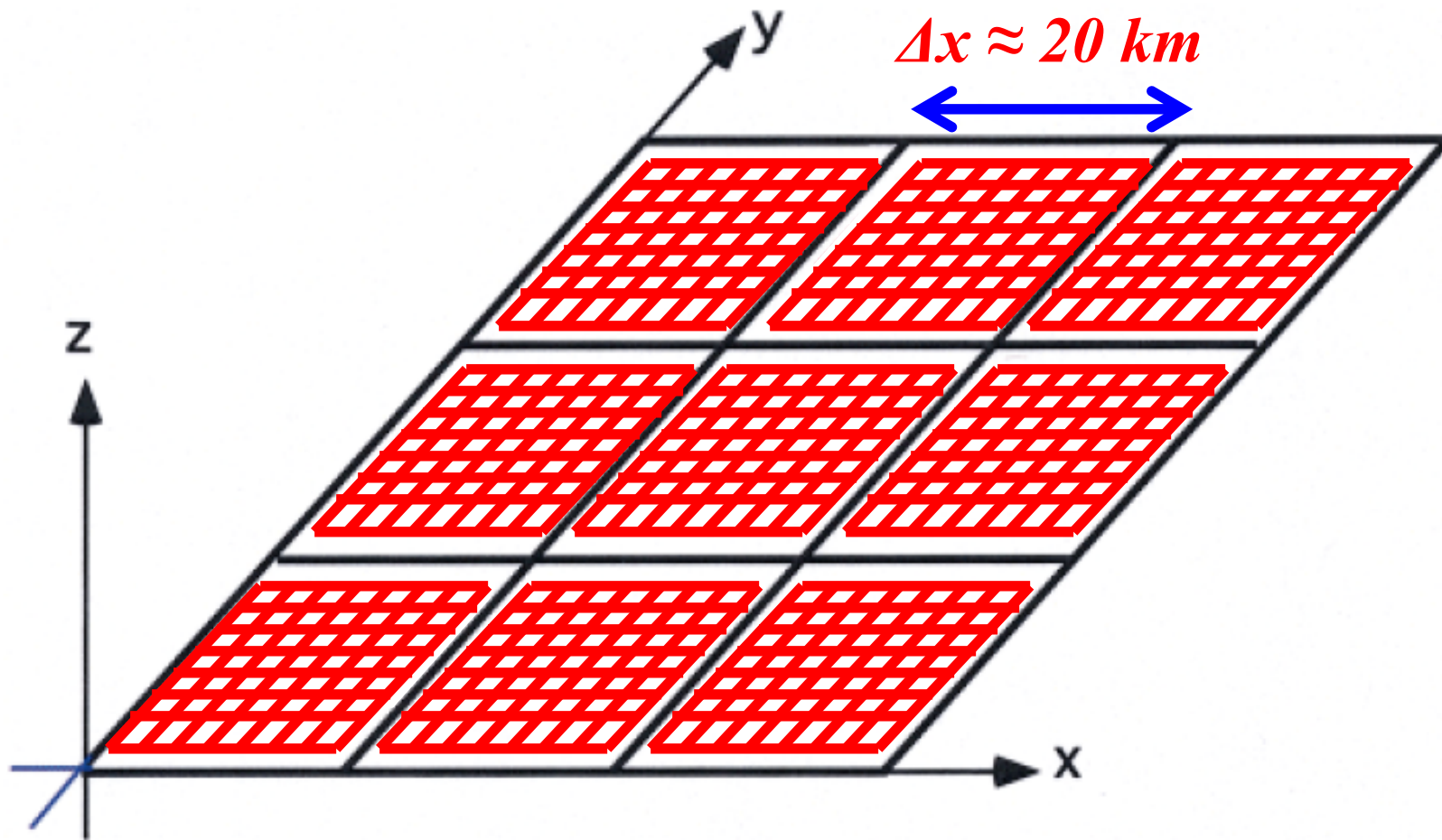
- Parallel processing?
- What equations to use?

SP can help! And can also provide additional benefits...

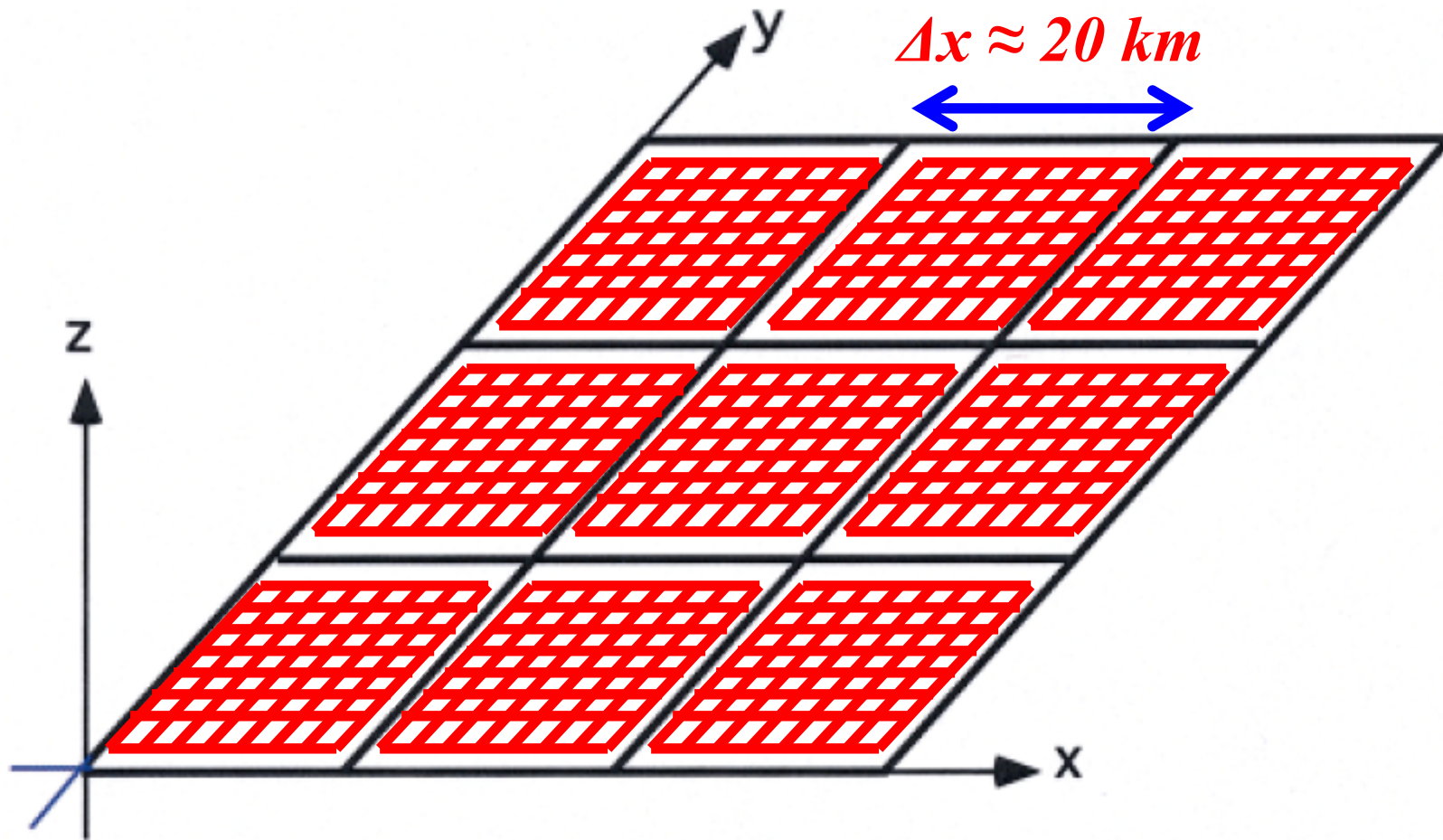
Original SP proposal:



Next generation SP proposal:



Next generation SP proposal:



Communication between the outer model and SP models takes place only through the profiles, see Grabowski (*JAS* 2004)

Issues:

- Parallel processing?

Not a problem! SP is embarrassingly parallel with small amount of data that needs to be transfer infrequently between the host model and SP 3D models (only the profiles). Ideal for GPUs!

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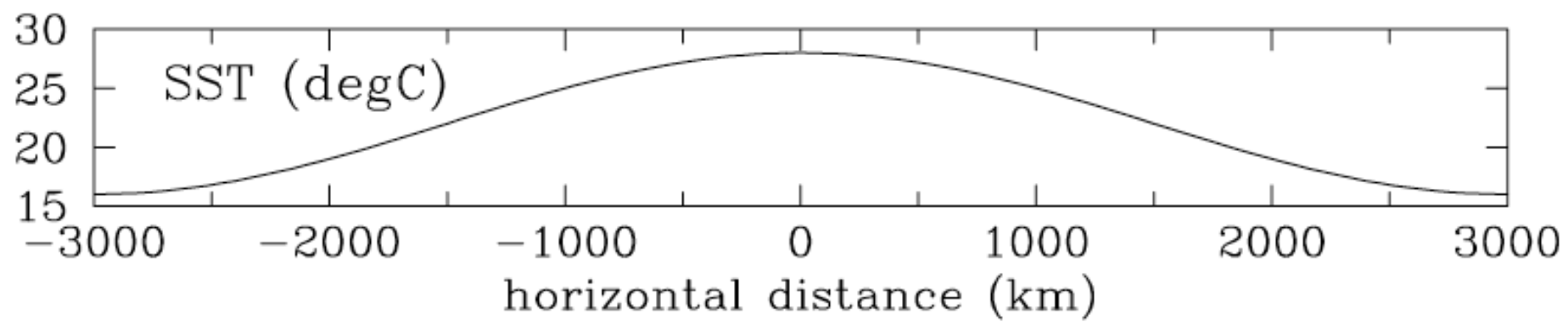
- SP can provide additional benefits:

SP models can have different grids, essentially allowing unstructured grid system with no additional cost.

Illustration: the 2D mock-Hadley circulation

Similar to mock-Walker circulation (Grabowski *JAS* 2000) but with a larger SST difference between ascending and descending branches (4 degC in mock-Walker versus 12 degC in mock-Hadley)

One expects deep convection over warm SSTs and stratocumulus-topped boundary layer over cold SSTs...



Model setup:

6,000 km horizontal domain

24 km vertical extent, with stretched grid

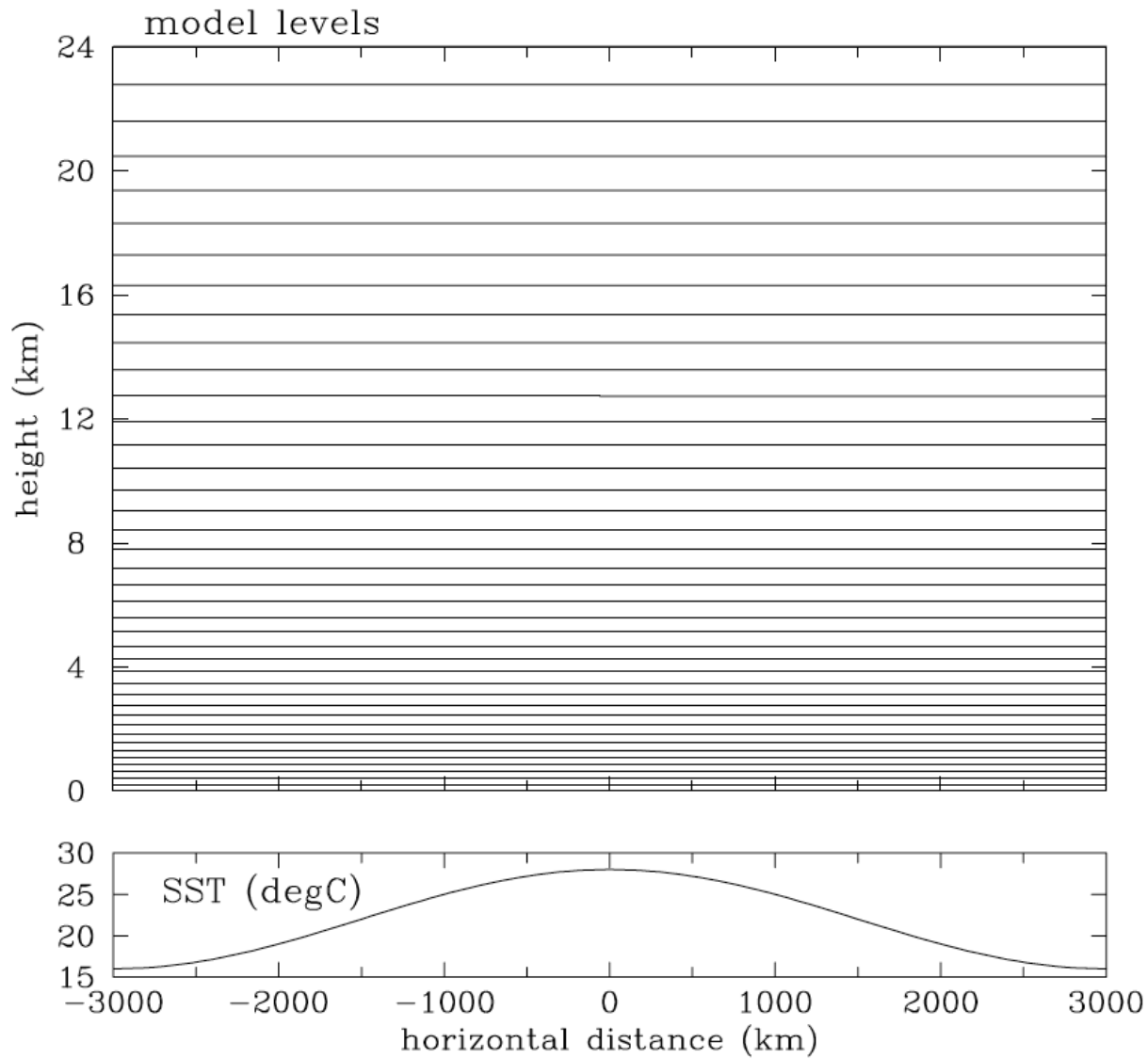
SST: 16 to 28 degC, varying as $\cos(\text{distance})$

No mean flow

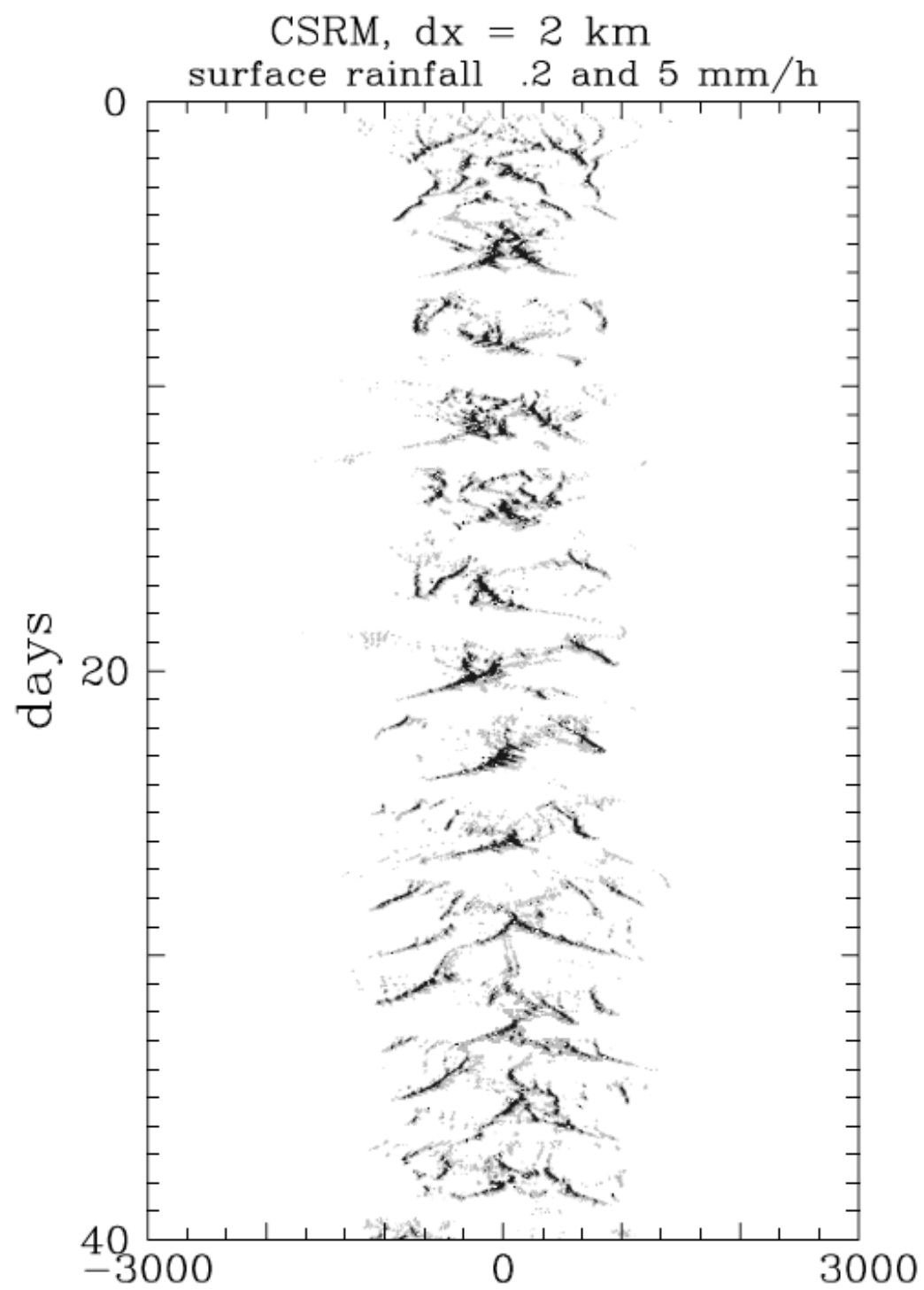
Prescribed radiative cooling: 1.5 K/day below 12 km,
decreasing linearly to zero at 15km

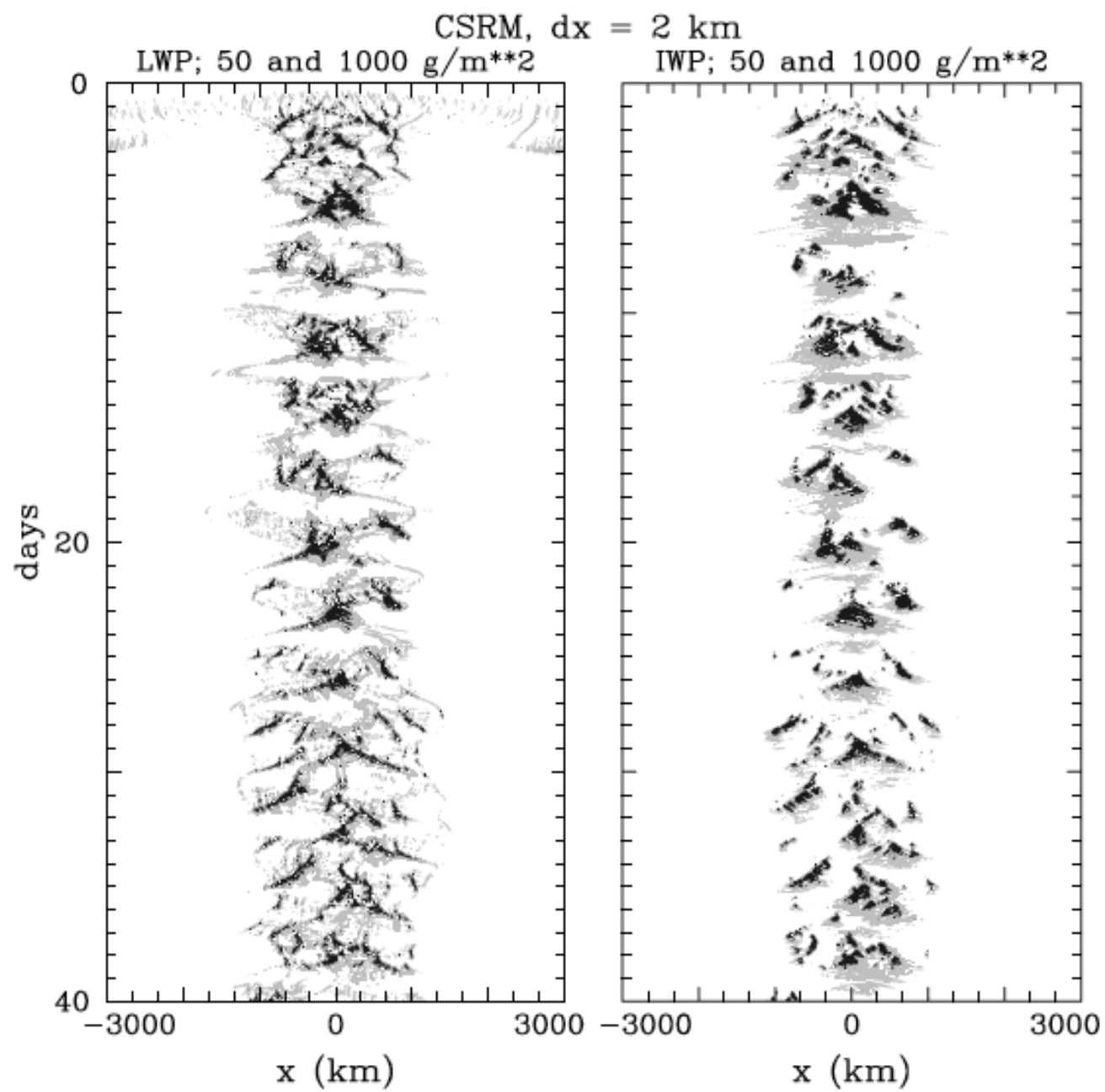
No SGS model in either outer or SP models (implicit LES)

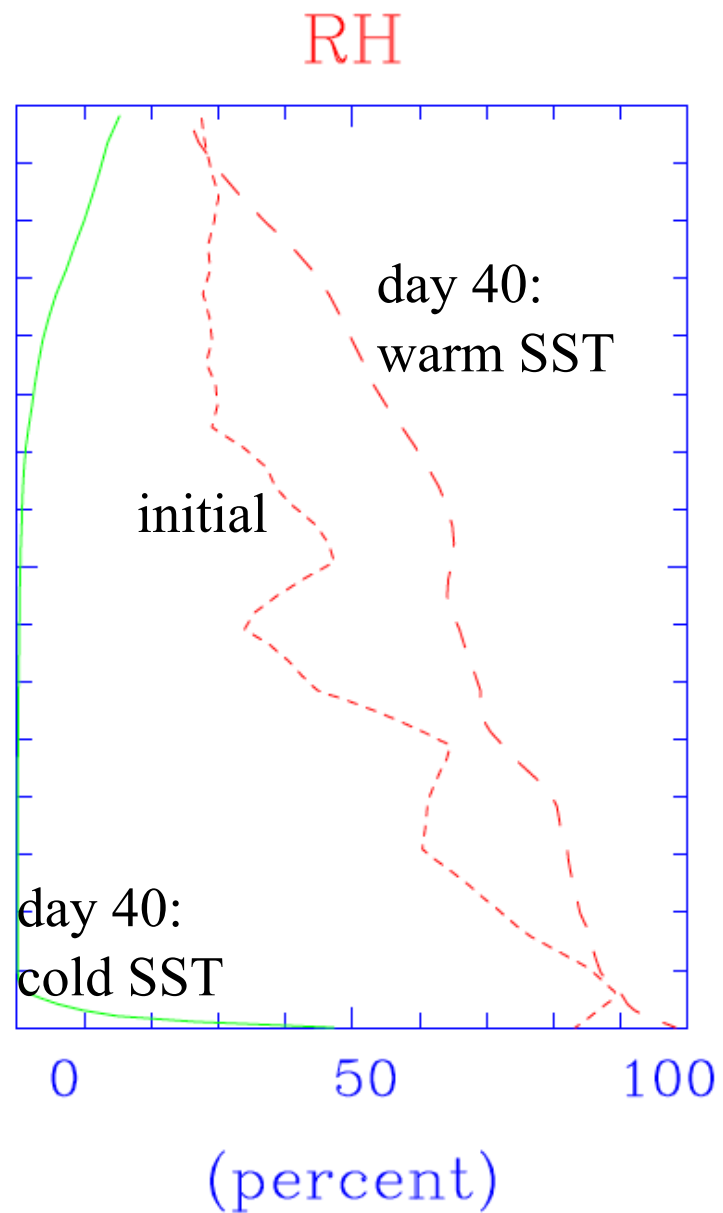
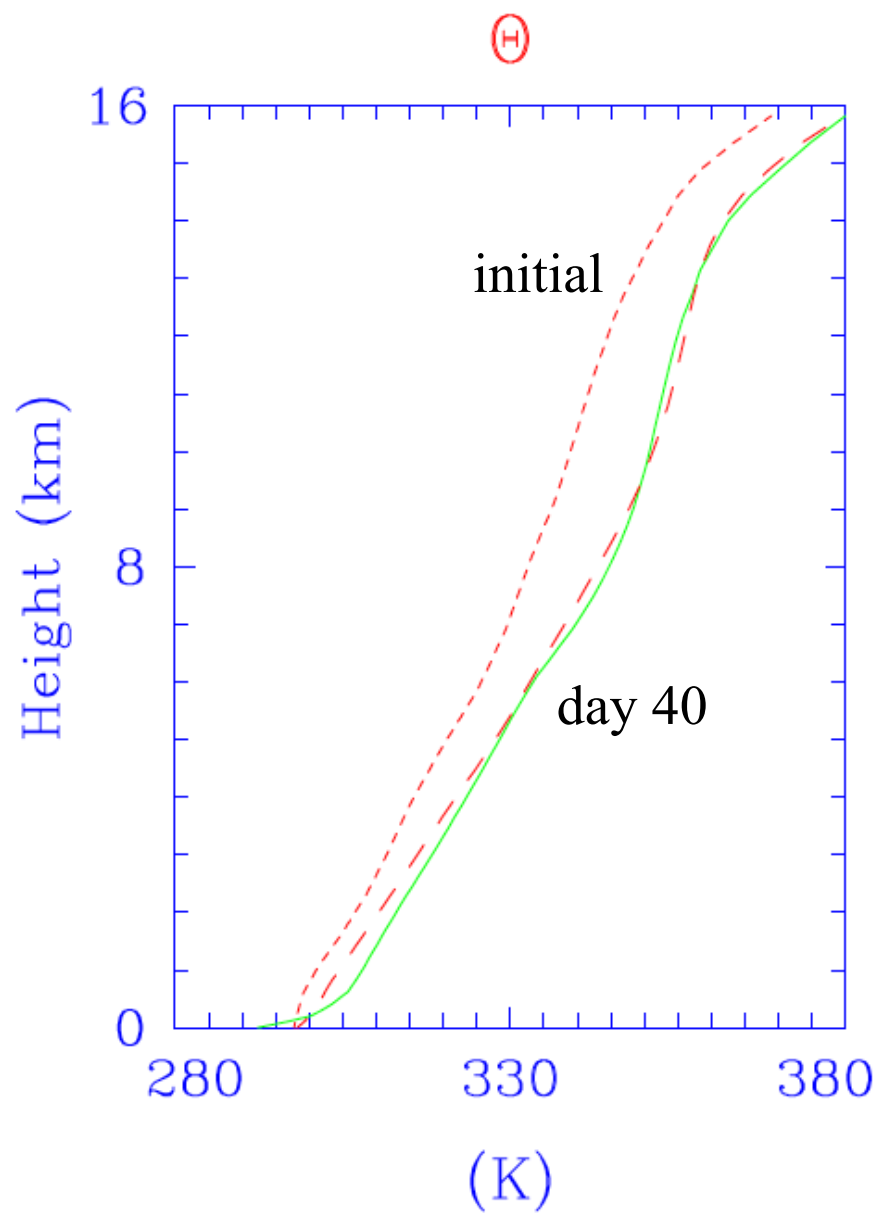
Simple formulation of surface sensible and latent heat fluxes

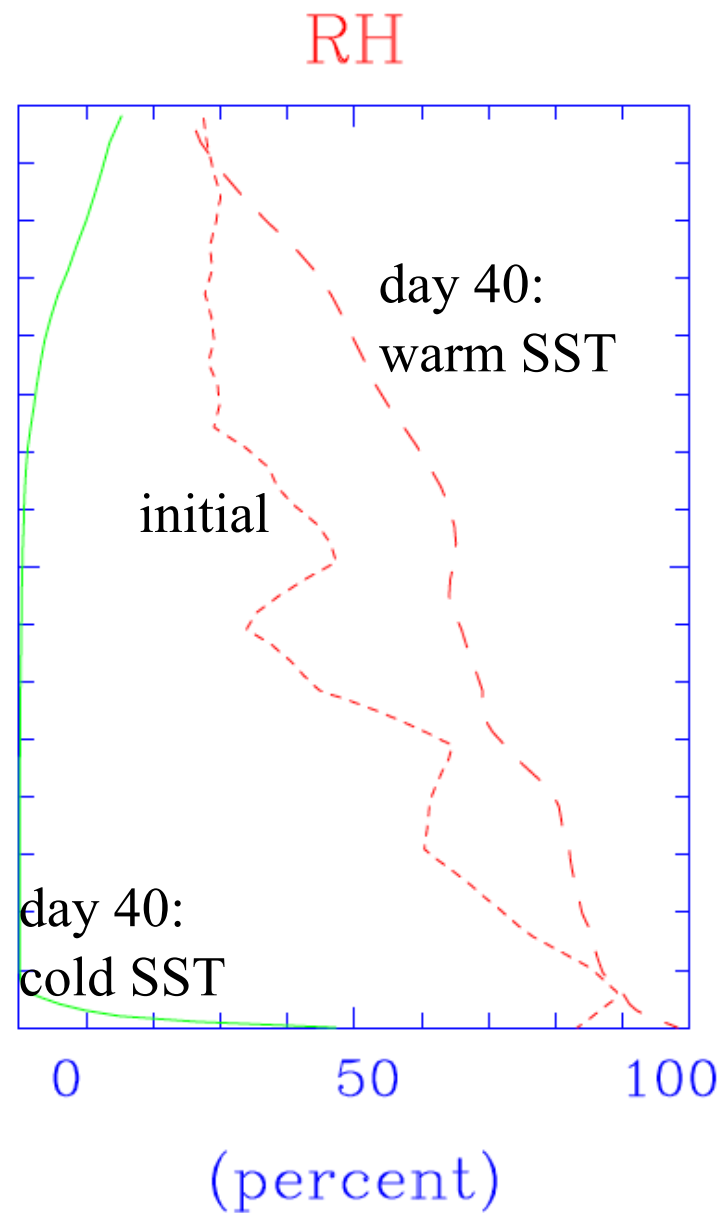
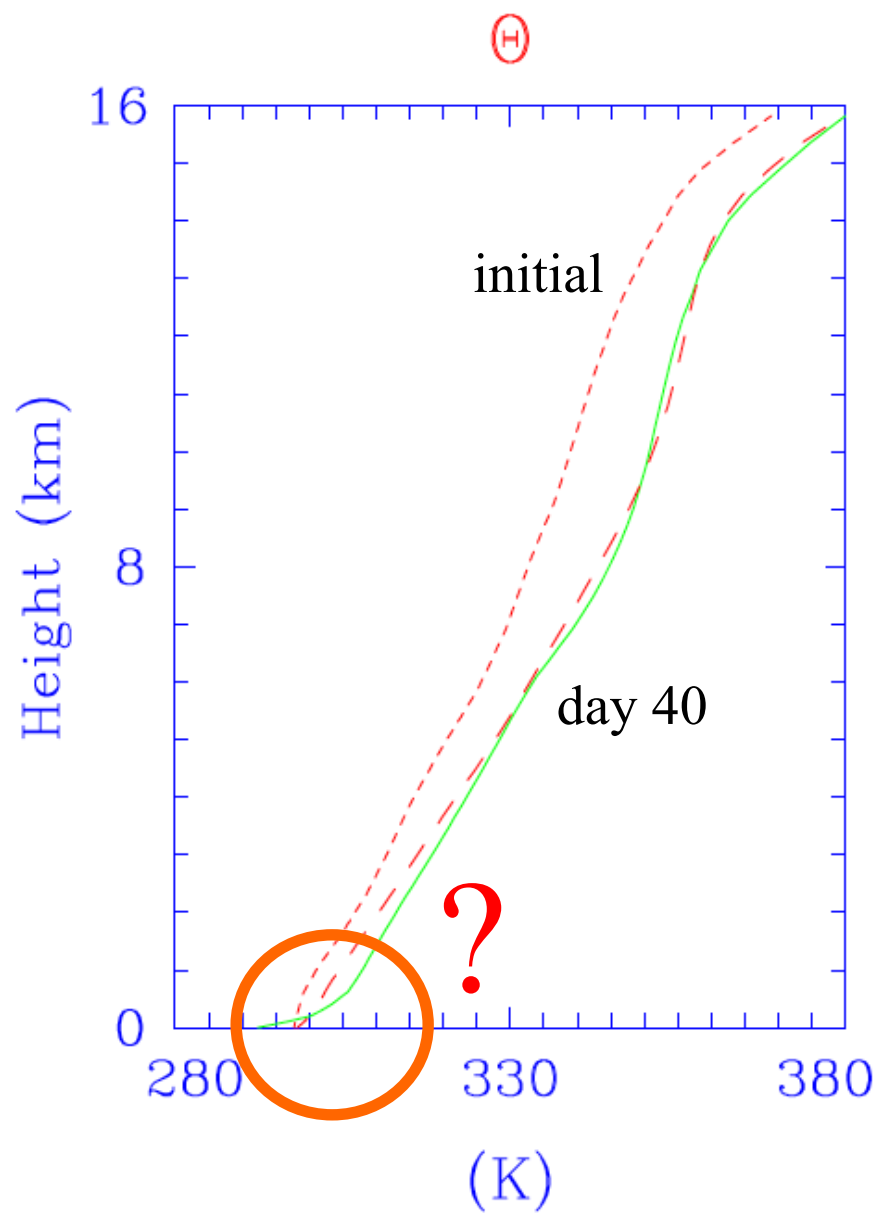


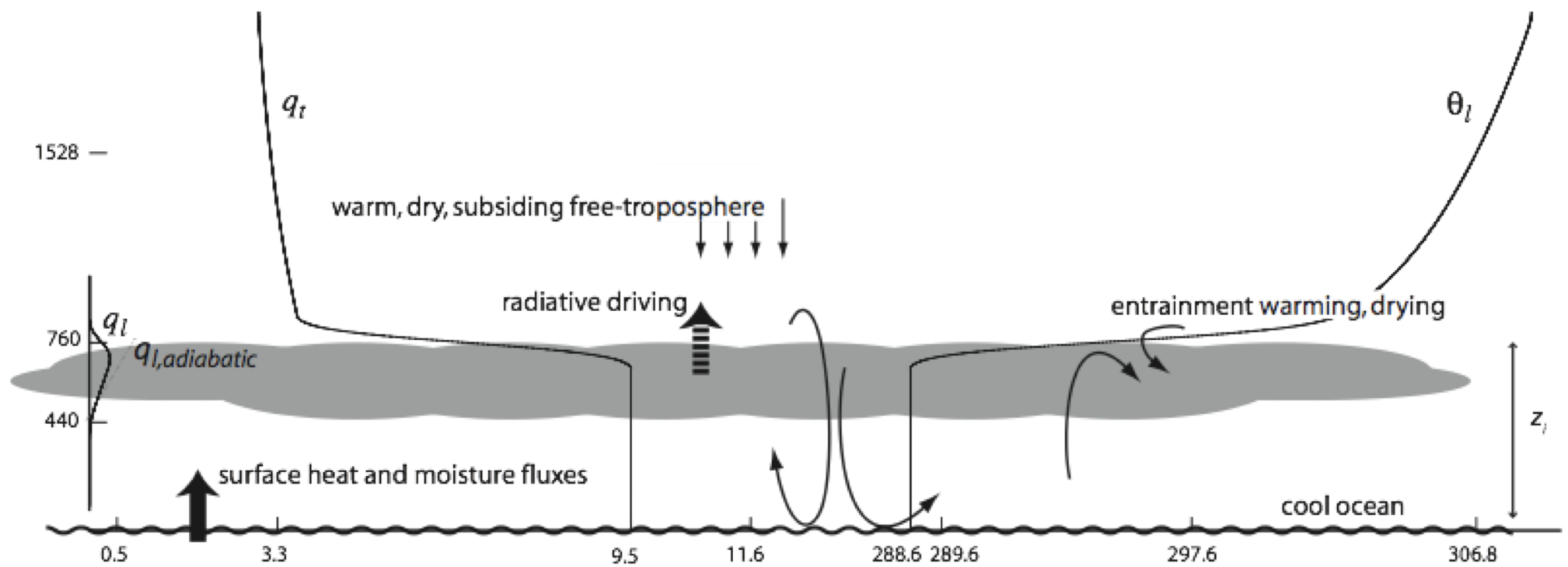
*Horizontal domain and
vertical grid for CRM
simulation, $\Delta x=2$ km,
3000 points in the
horizontal, 81 levels.*



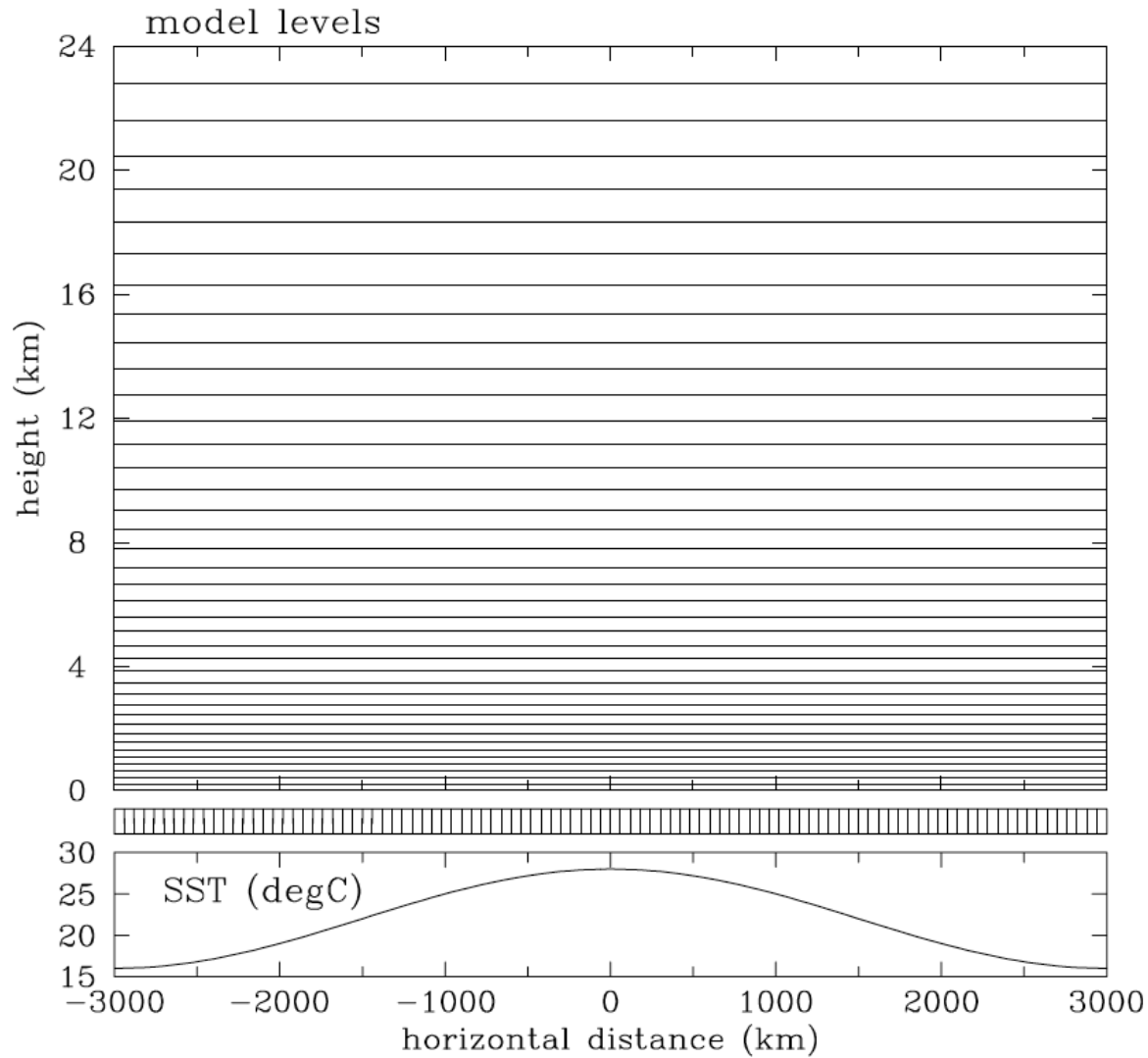








Stevens et al., 2006 , MWR

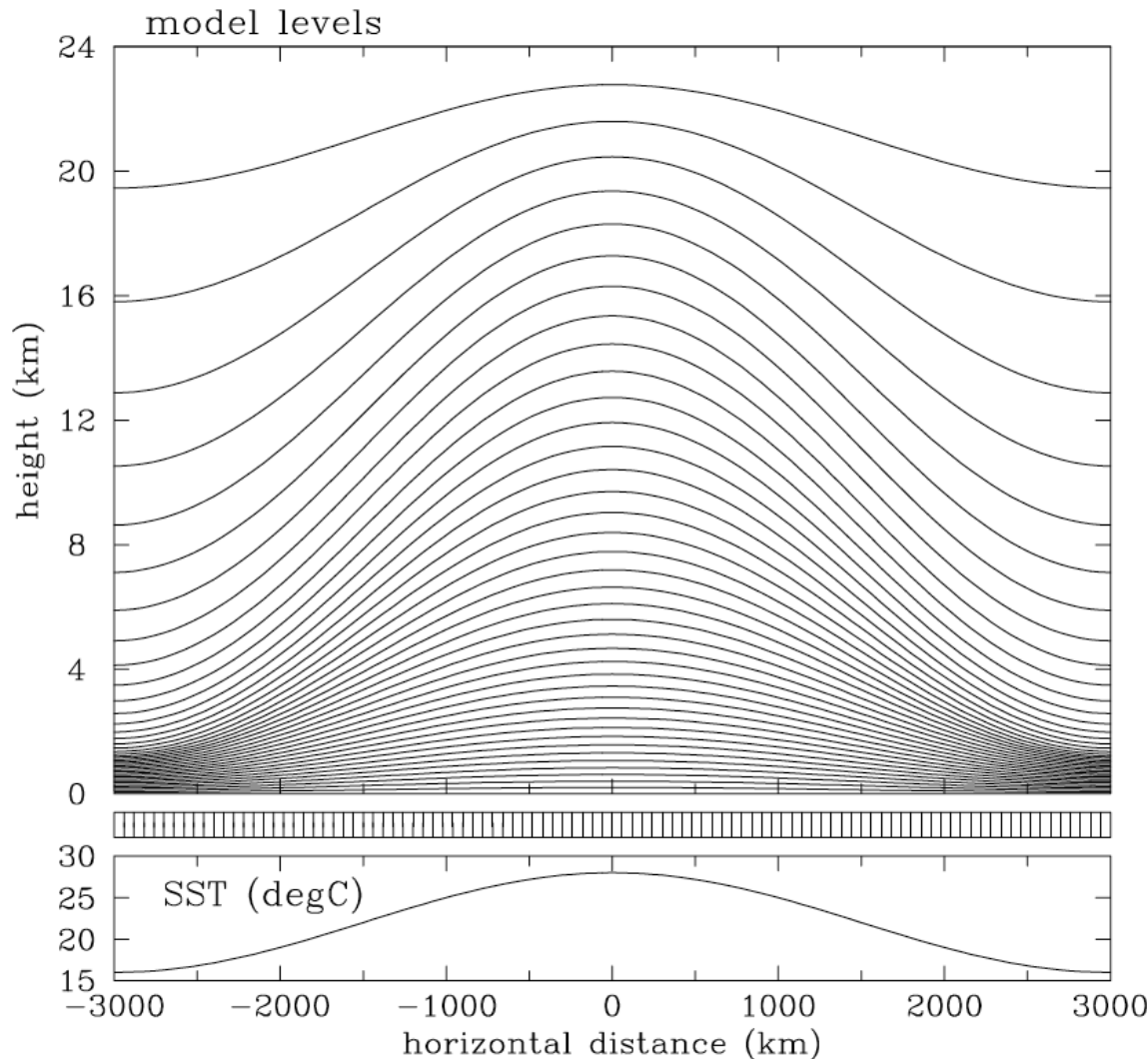


Traditional SP model:

*Outer model: $\Delta x=60$ km,
100 points in the
horizontal, 81 levels.*

*SP models: $\Delta x=2$ km, the
same vertical grid as the
outer model.*

Heterogeneous SP model:



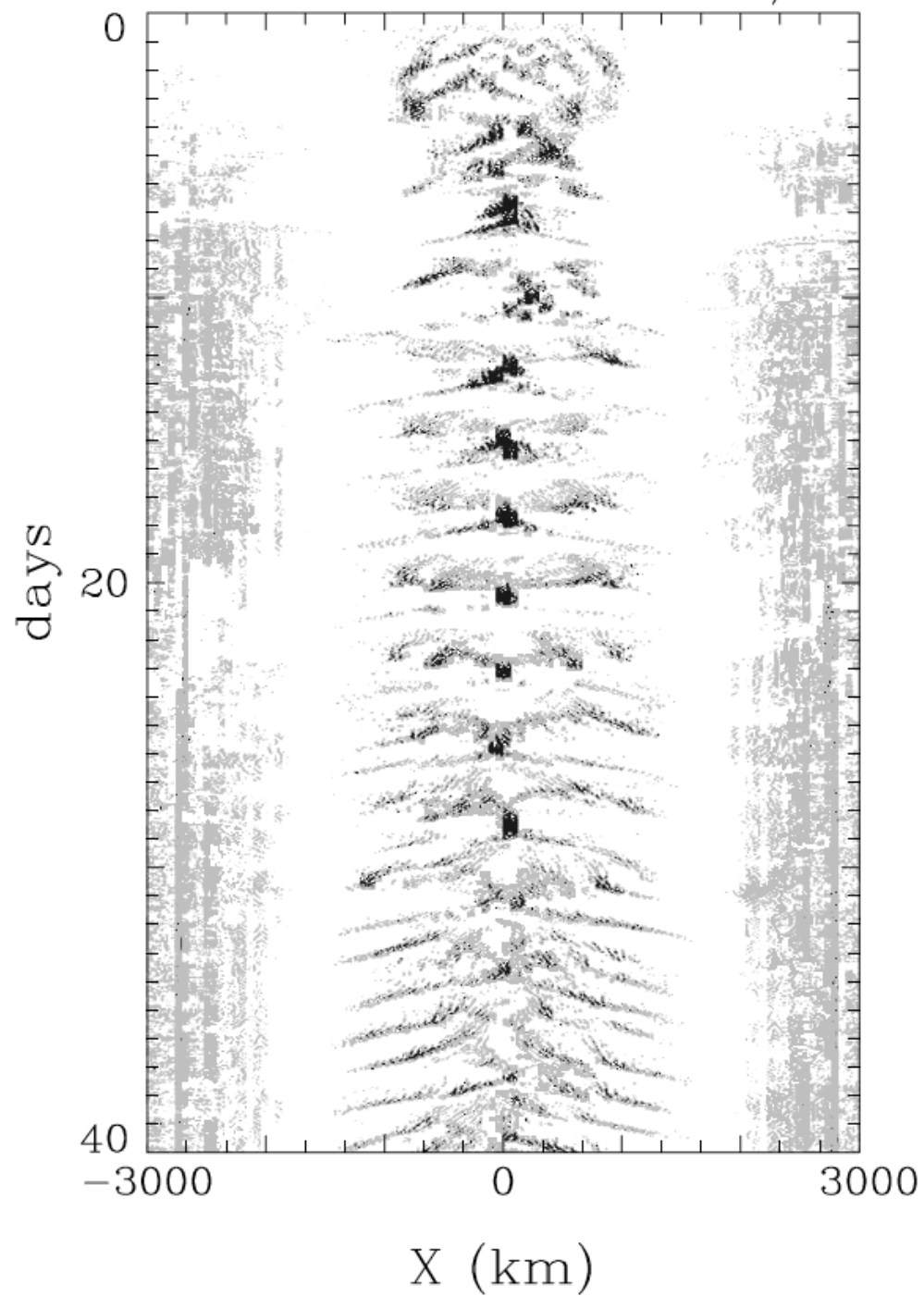
Outer model: $\Delta x=60$ km, 100 points in the horizontal, 81 levels.

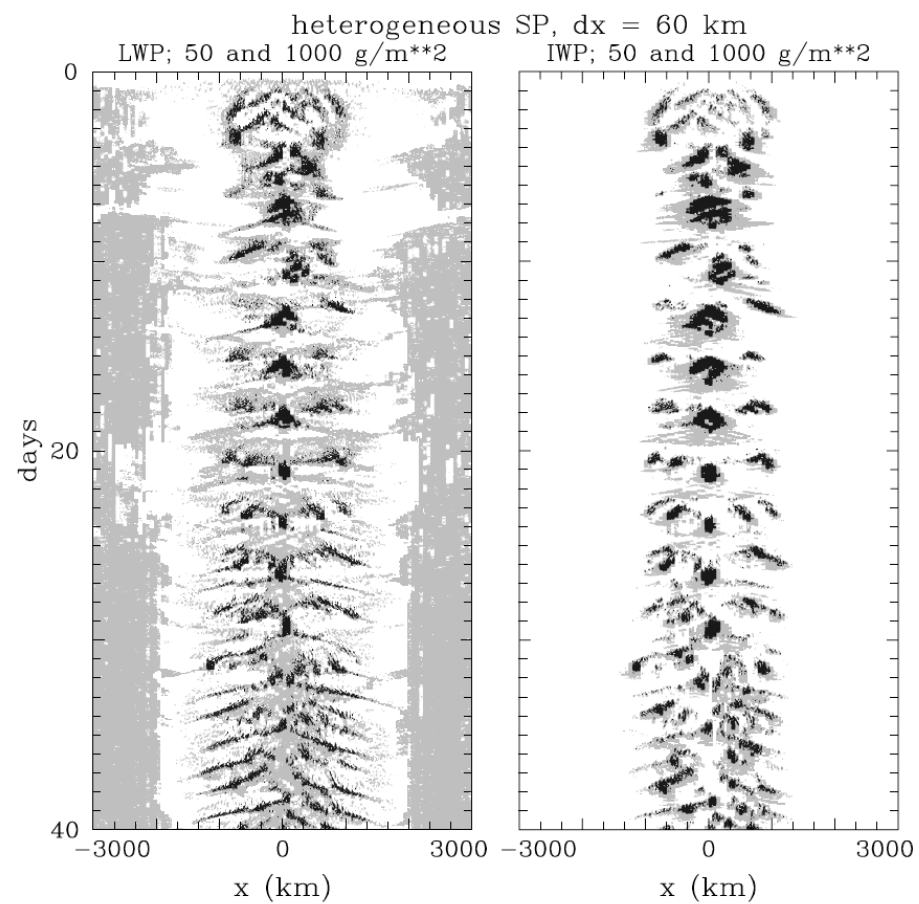
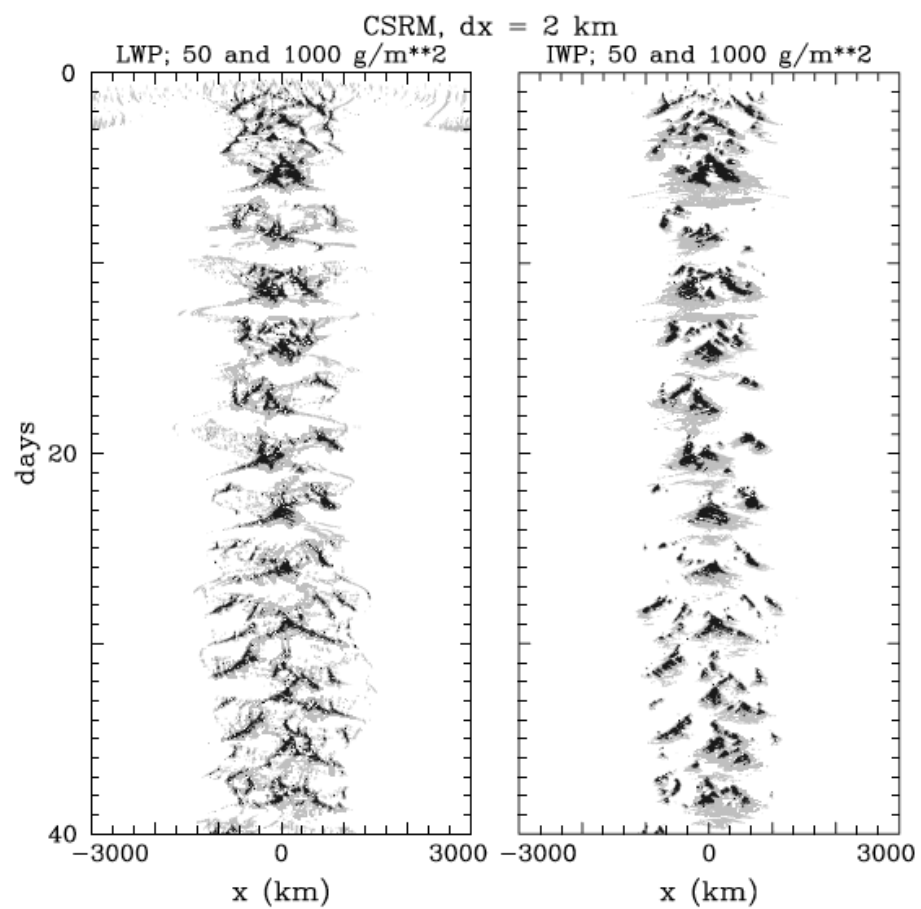
*SP models at high SST:
CRM: $\Delta x=2$ km, the same vertical grid as the outer model.*

*SP models at low SST:
“2D LES”: $\Delta x=200$ m, stretched vertical grid with $\Delta z=30$ m below 1 km, stretching strongly above.*

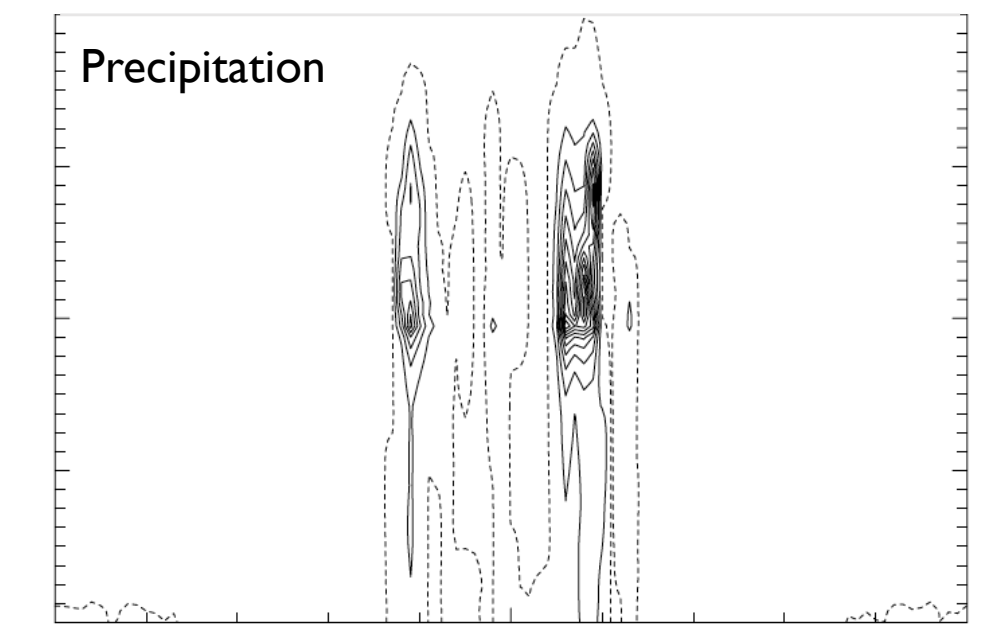
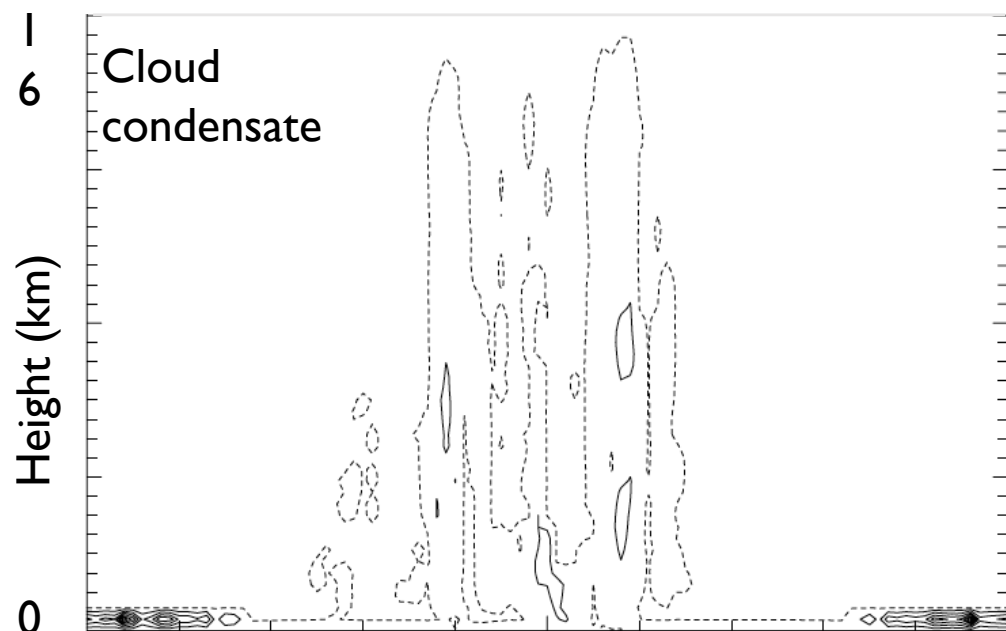
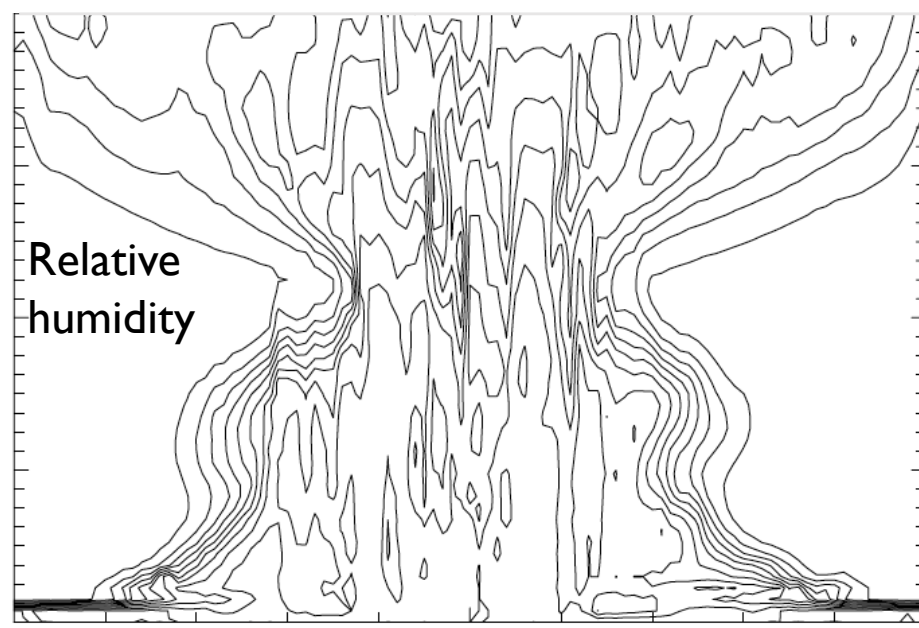
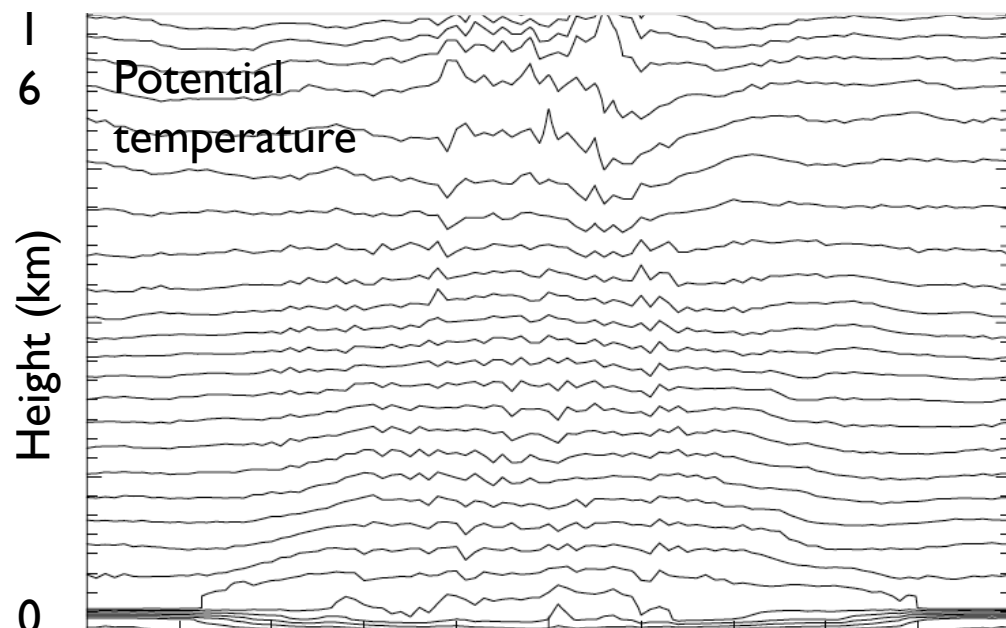
Linear interpolation of profiles between outer and SP models.

heterogeneous SP, $dx = 60$ km
surface rainfall .2 and 5 mm/h





Snapshots of fields at day 40 as seen on the outer model grid...



-3000

0

3000

Distance (km)

-3000

3000

Distance (km)

Conclusions:

1. *Large eddy simulation (LES)* provides an appropriate framework for modeling cloud processes in both shallow boundary layer clouds and deep convection. The race towards *global LES* is on.

2. A brute force approach, that is, extending global convection-permitting models (such as the Japanese NICAM or German ICON) to global LES will be computationally extremely expensive because of the amount of data that needs to be transferred between subdomains in traditional parallelization methodologies. The efficiency of the compressible dynamical framework at such resolutions is also unclear.

Conclusions, continued:

3. The super-parameterization (SP) methodology provides a rapid way forward towards global LES. Outer model should have tiles of 100s km² (say 20 by 20 km) and can be hydrostatic. 3D SP models can be anelastic with base-state and environmental profiles varying between equator and the poles, and they can have different grids depending on geographic location. Parallelization of such a system is trivial with only profiles exchanged infrequently between outer and SP models. Such a global LES system based on SP methodology should run efficiently on massively parallel systems, for instance, those based on GPUs.