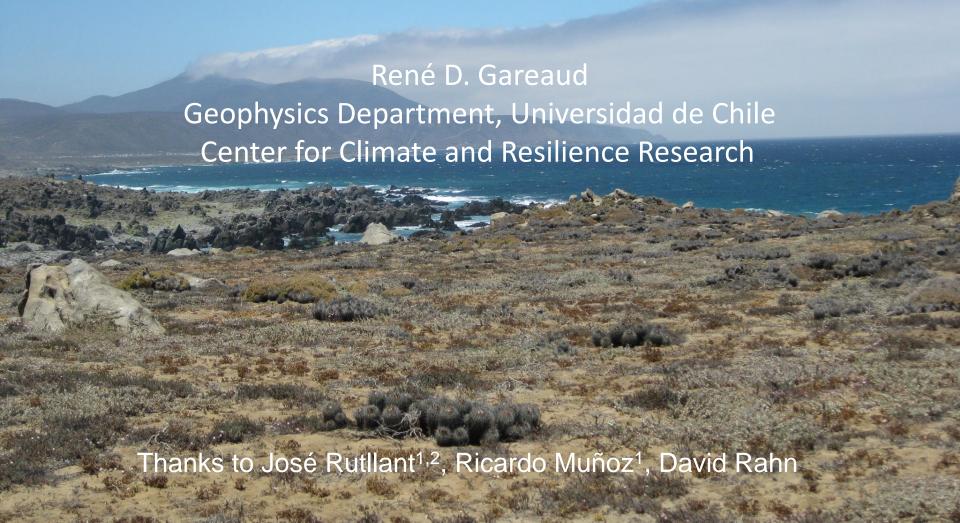
# Cloud over EBUS Impact and response



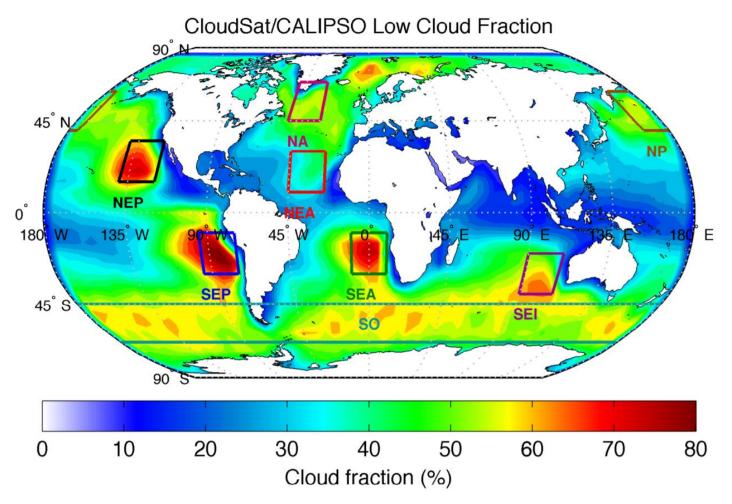
# Outline

Our cloudy EBUS...some mesoscale aspects

Large scale impact

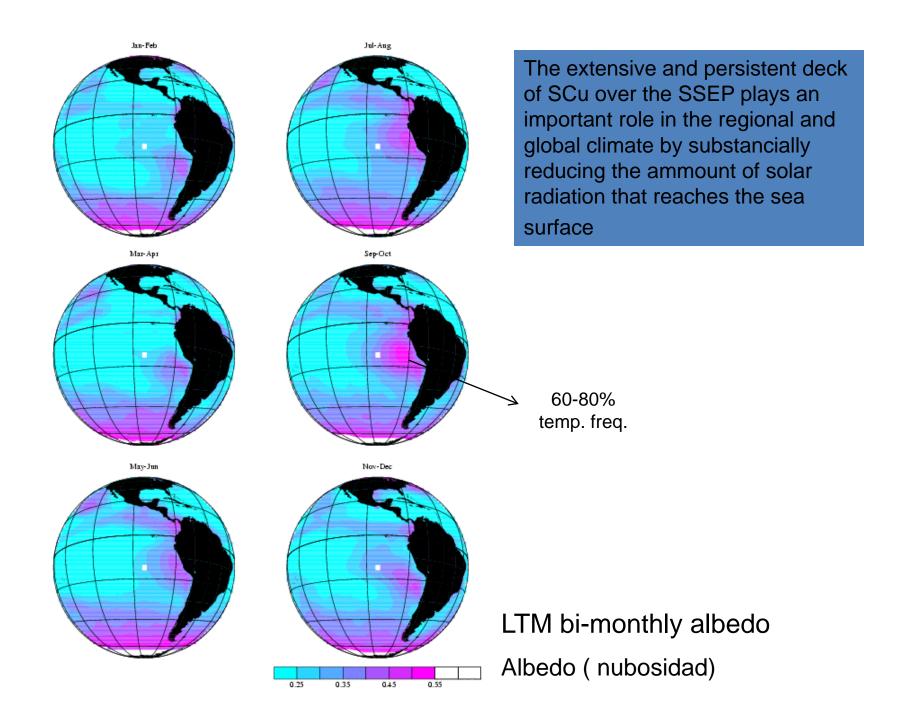
Challenges

## EBUS: Cold SST and Cloudy Regions

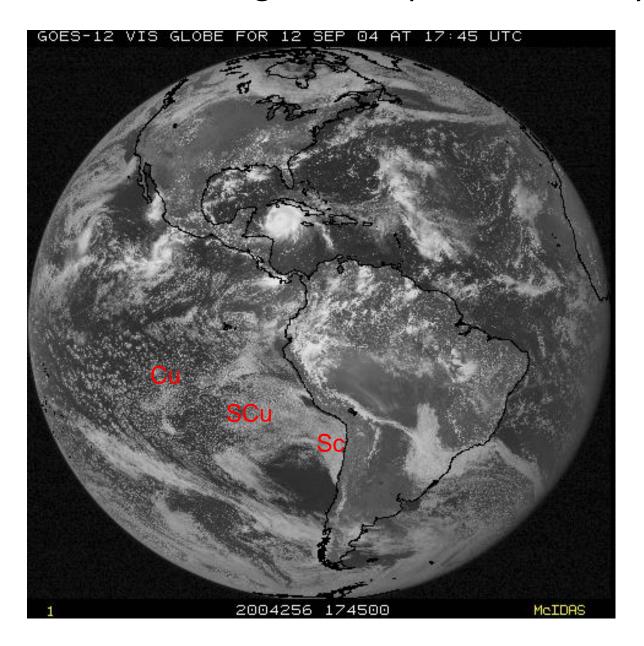


Climatology of stratocumulus cloud morphologies: microphysical properties and radiative effects

A. Muhlbauer<sup>1</sup>, I. L. McCoy<sup>2</sup>, and R. Wood<sup>3</sup>



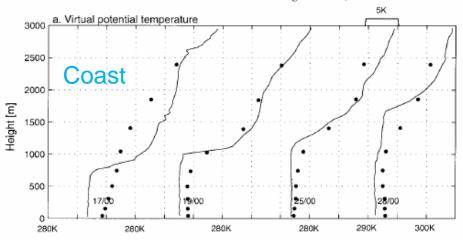
## Cloud field has significant spatial variability



#### CIMAR-5: A Snapshot of the Lower Troposphere over the Subtropical Southeast Pacific



René D. Garreaud,\* José Rutllant,\* Juan Quintana,+ Jorge Carrasco,+ and Patrick Minnis\*



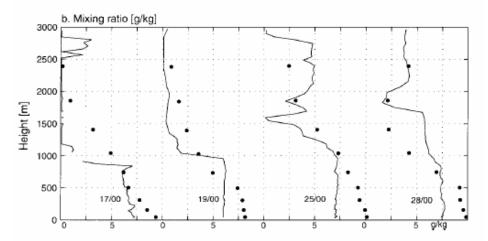
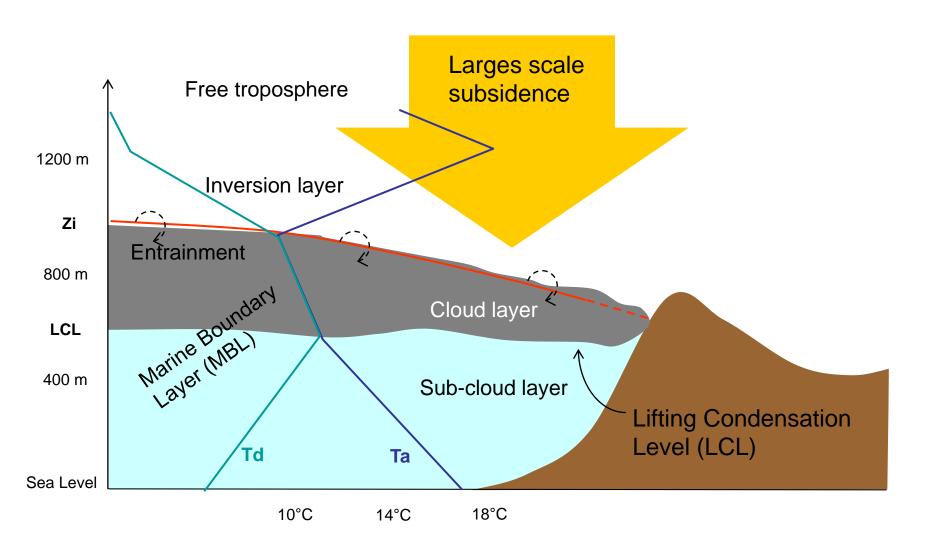


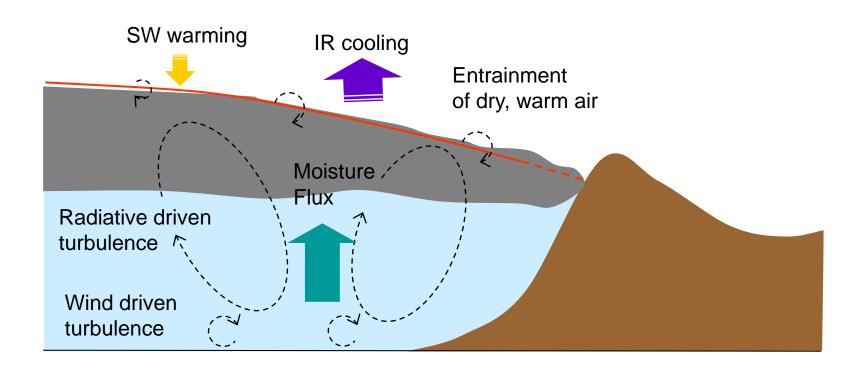
Fig. 7. (a) Vertical profiles of virtual potential temperature ( $\theta_v$ ) at four times during CIMAR-5. Sounding time is indicated by day (Oct)/UTC at the side of each profile. Dots are values of  $\theta_v$  from the NCEP–NCAR reanalysis (original  $\sigma$  levels) interpolated to the ship position. (b) As in (a) but for mixing ratio.

The inversion layer is maintained by the large-scale subsidence. It caps a cool, moist, often well mixed marine boundary layer in contact with the cold waters of the SE Pacific.

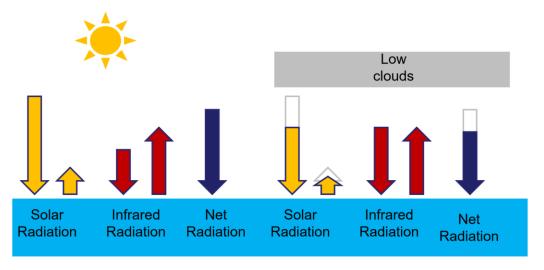


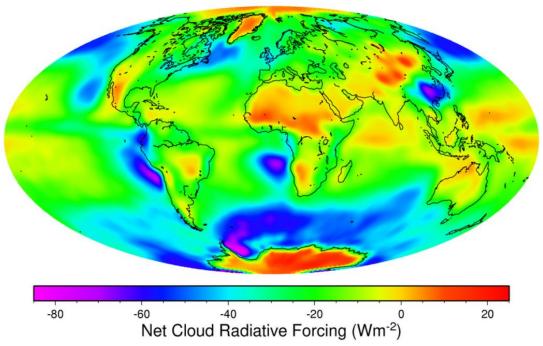
The turbulence within the MBL is largely driven by longwave cooling at the top of the cloud deck. Large eddies transport moist air upward, eventually reaching the LCL and forming the cloud. Note the feedback between cloud and turbulence.

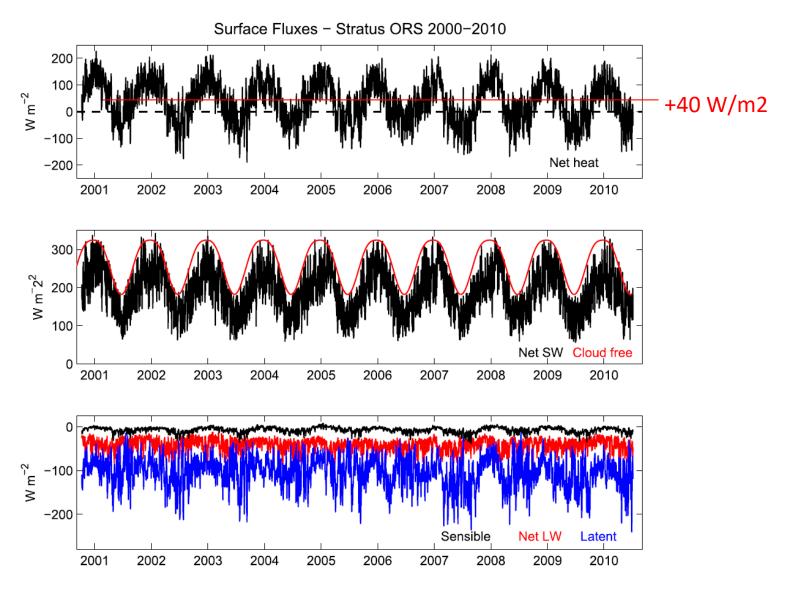
Eddies can also overshoot the MBL thus entraining dry, warm air that tend to disipate the cloud.



# EBUS: Cold SST and Cloudy Regions

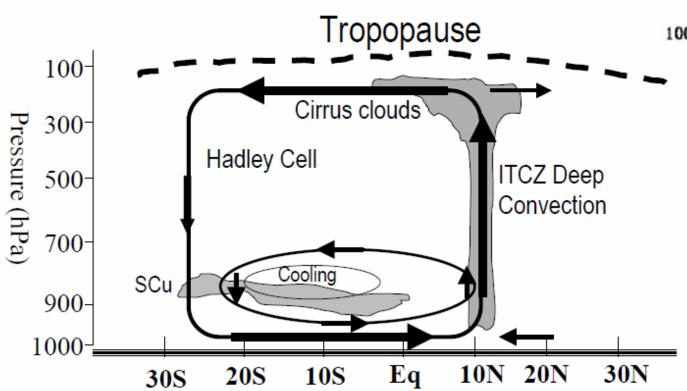


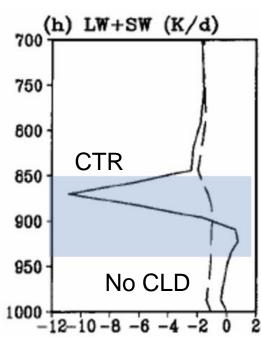




#### Large-Scale Atmospheric Forcing by Southeast Pacific Boundary Layer Clouds: A Regional Model Study\*

YUQING WANG, SHANG-PING XIE, + BIN WANG, + AND HAIMING XU





Wang et al. 2004

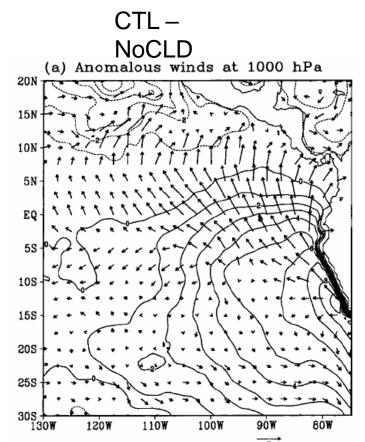
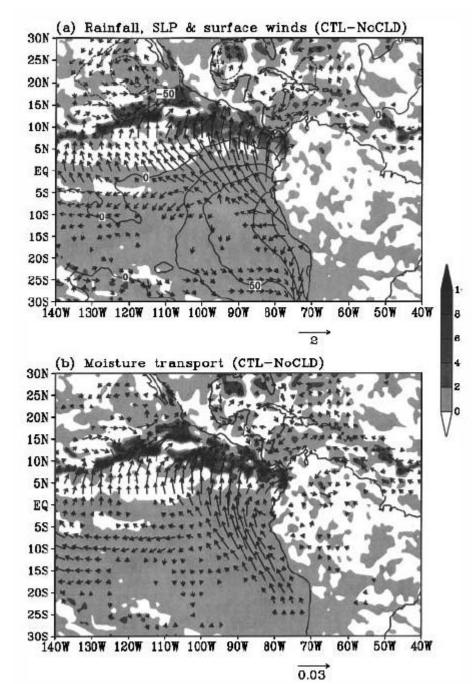
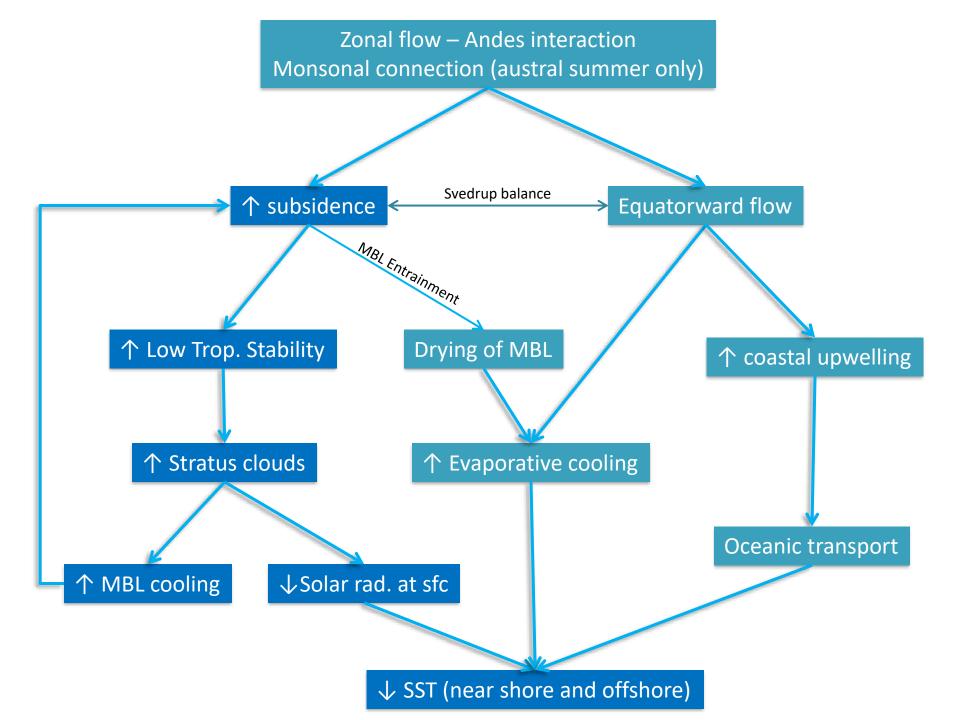


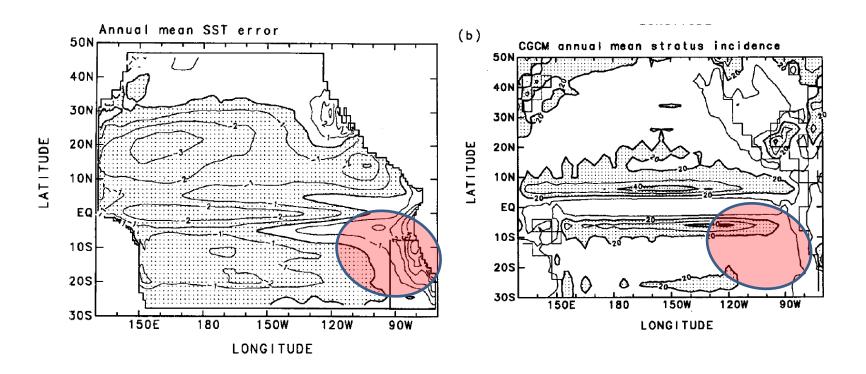
Fig. 9. (a) Differences in ASO mean daily rainfall (shaded in mm day $^{-1}$ ), sea level pressure (contours in Pa), and the 10-mheight winds (vectors with unit vector of 2 m s $^{-1}$ ) between CTL and NoCLD (CTL-NoCLD); (b) difference in column-integrated moisture transport (vectors with unit vector of 0.03 kg kg $^{-1}$  m s $^{-1}$ ) between CTL and NoCLD superposed onto difference in rainfall as in (a).





# Control (original) UCLA GCM simulation

# few clouds, large warm bias



Peruvian Stratus Clouds and the Tropical Pacific Circulation: A Coupled Ocean-Atmosphere GCM Study

CHUNG-CHUN MA, CARLOS R. MECHOSO, ANDREW W. ROBERTSON, AND AKIO ARAKAWA

#### SST differences: Fixed SCu-Control

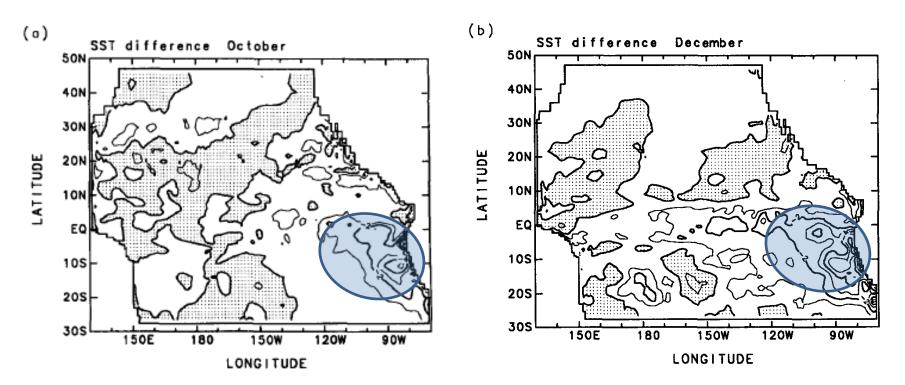


Fig. 7. Difference in simulated SST between the stratus experiment and the control simulation for (a) October and (b) December. Contour interval is 1 K; regions with positive values are stippled.

Peruvian Stratus Clouds and the Tropical Pacific Circulation: A Coupled Ocean-Atmosphere GCM Study

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1644 JOURNAL OF CLIMATE VOLUME 9

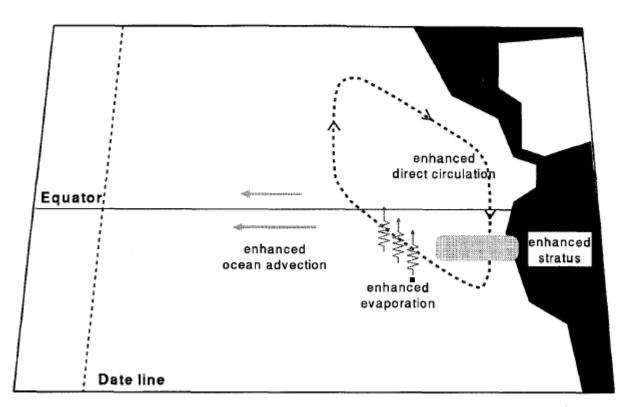
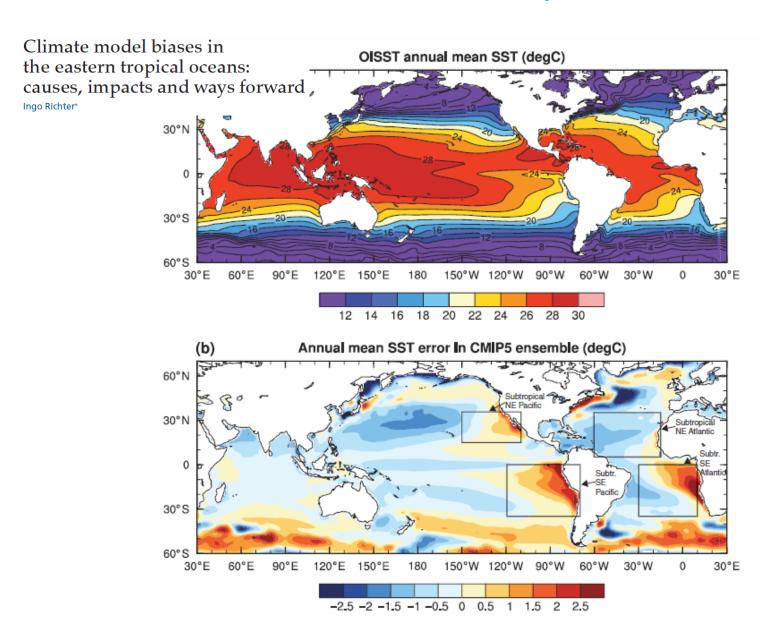


FIG. 10. A schematic illustrating the influence of Peruvian stratus on tropical atmospheric and oceanic circulations as suggested by the coupled GCM experiment.

### Huston....we have a problem



# Cloud – SST @ EBUS

#### Chicken and egg problem?

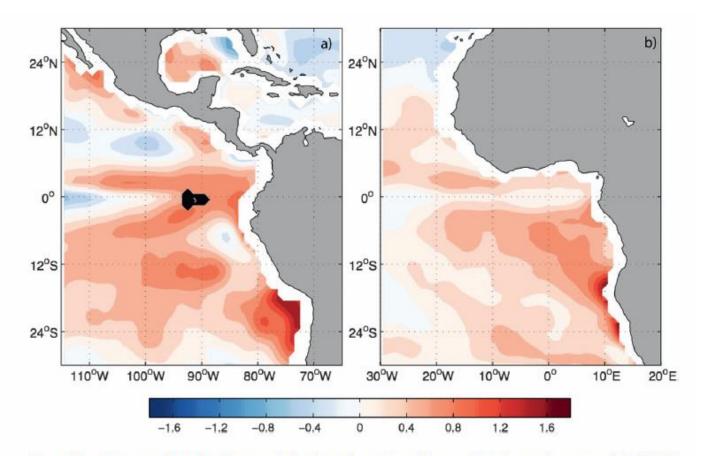
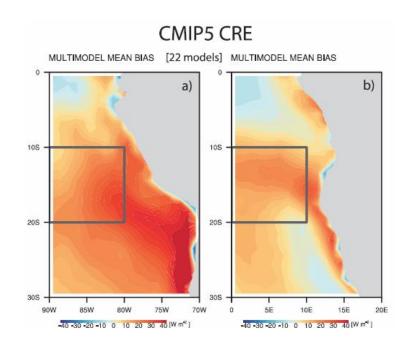


Fig. 10. Ocean simulations with fixed atmosphere forcings (termed OMIP) also produce SST biases, if less pronounced than in CMIP simulations, as shown in the 22-ensemble OMIP SST bias relative to CORE2 surface forcing for (a) the Pacific and (b) the Atlantic (Danabasoglu et al. 2014). This suggests oceanic origins also contribute to the SST biases.

# Cloud – SST @ EBUS

Chicken and egg problem?

Fully coupled Ocean-Atmosphere Model: large cloud biases



# Cloud – SST @ EBUS

#### Chicken and egg problem?

Fully coupled Ocean-Atmosphere Model: large cloud biases

Not fully alleviated by prescribing SST. Low clouds are hard to simulate!

