

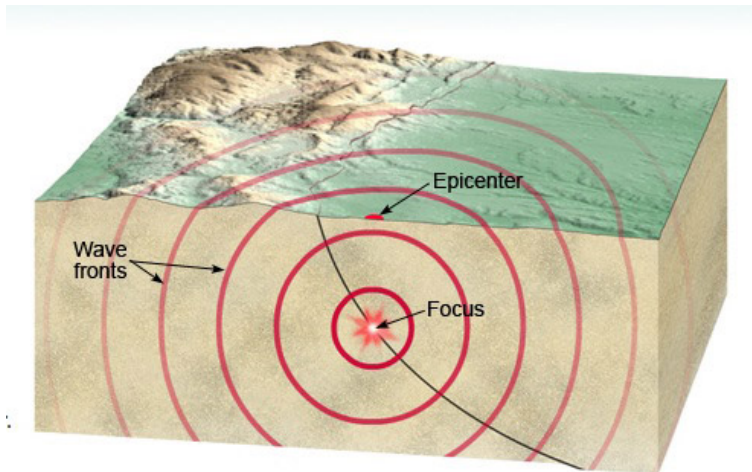
Advanced Workshop on Earthquake Fault Mechanics: Theory, Simulation and Observations

ICTP, Trieste, Sept 2-14 2019

Lecture 4: beyond friction

Jean Paul Ampuero (IRD/UCA Geoazur)

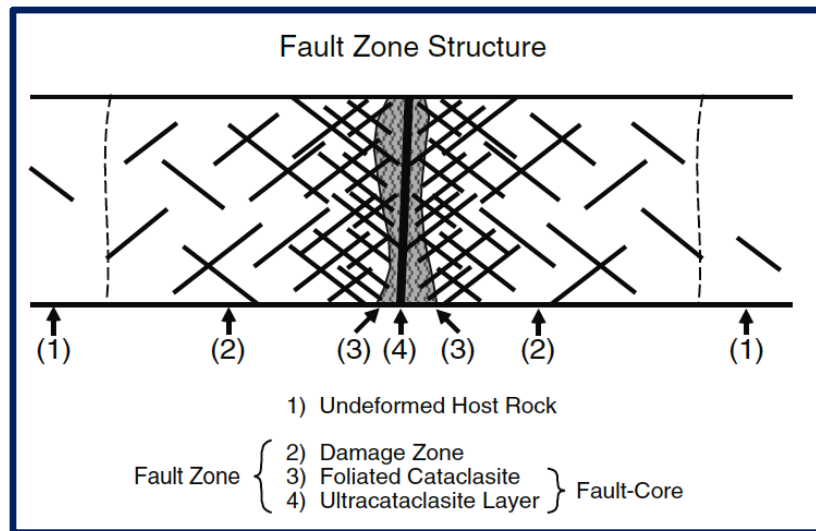
Real faults are thick ...



Punchbowl fault, CA
(Chester and Chester, 1998)

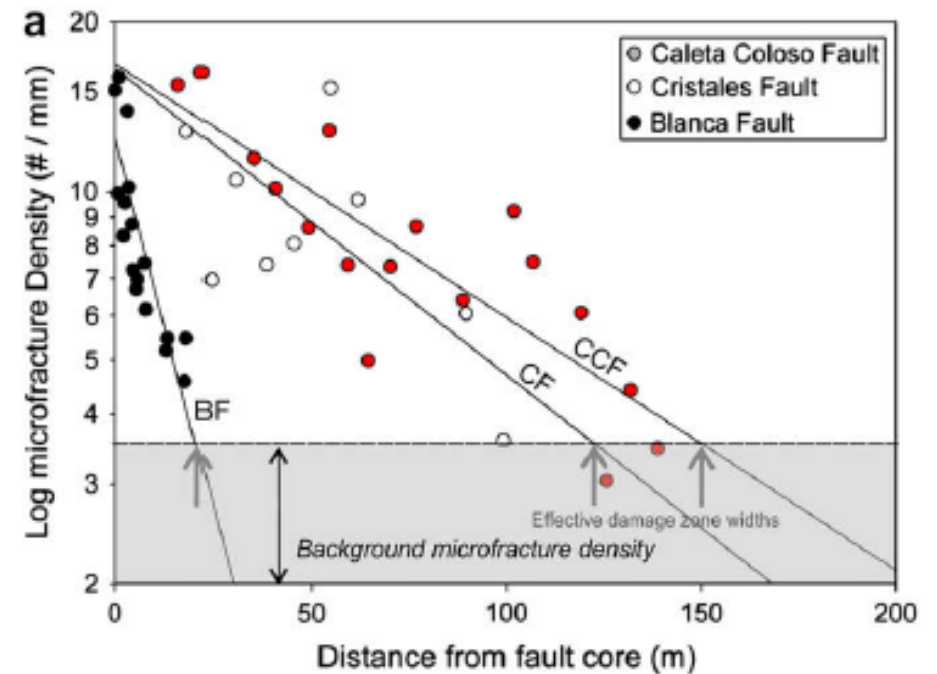


Fault zones are damaged

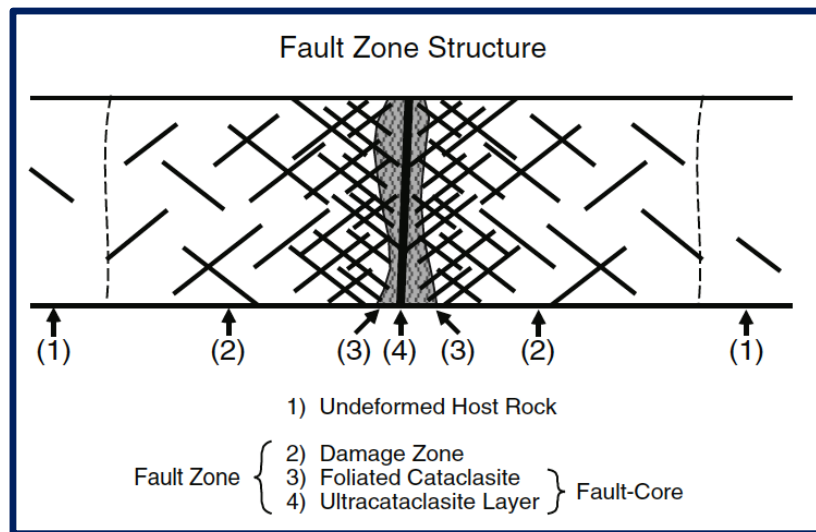


(Chester et al., 2004)

Mitchell and Faulkner (2009)

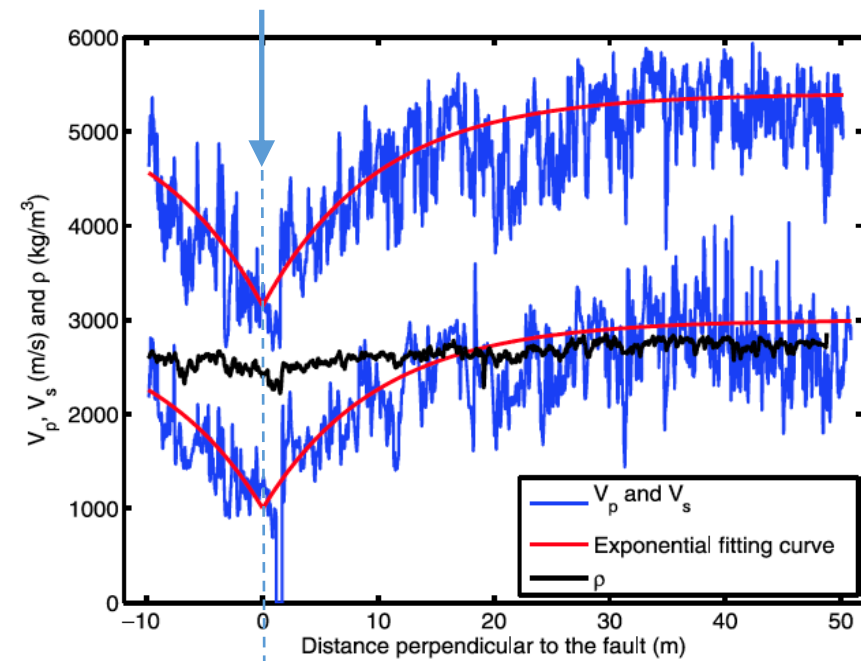


Fault zones are damaged



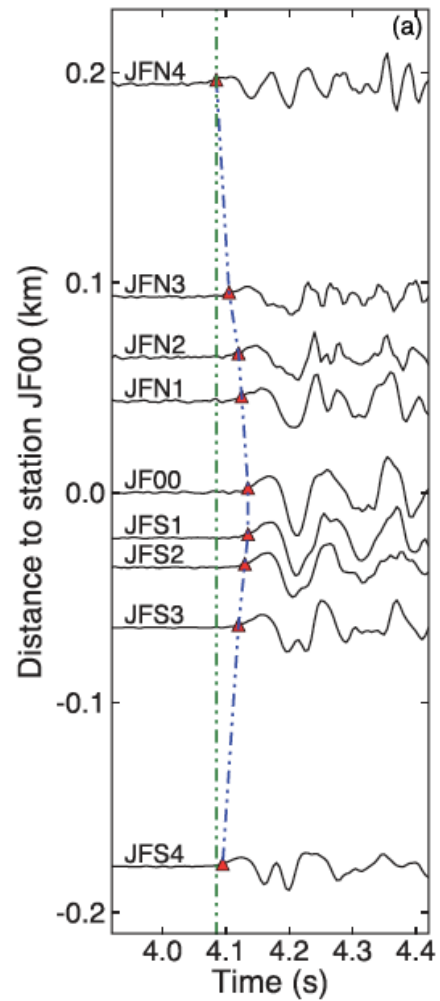
(Chester et al., 2004)

Low wave velocity zone surrounding the Nojima Fault, Kobe, Japan



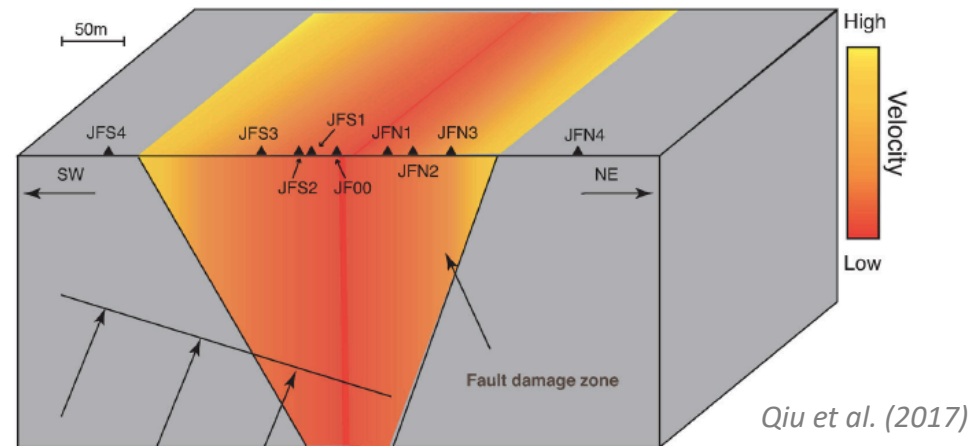
(Huang and Ampuero, 2011;
borehole data courtesy of H. Ito)

Low velocity fault zones imaged by trapped waves

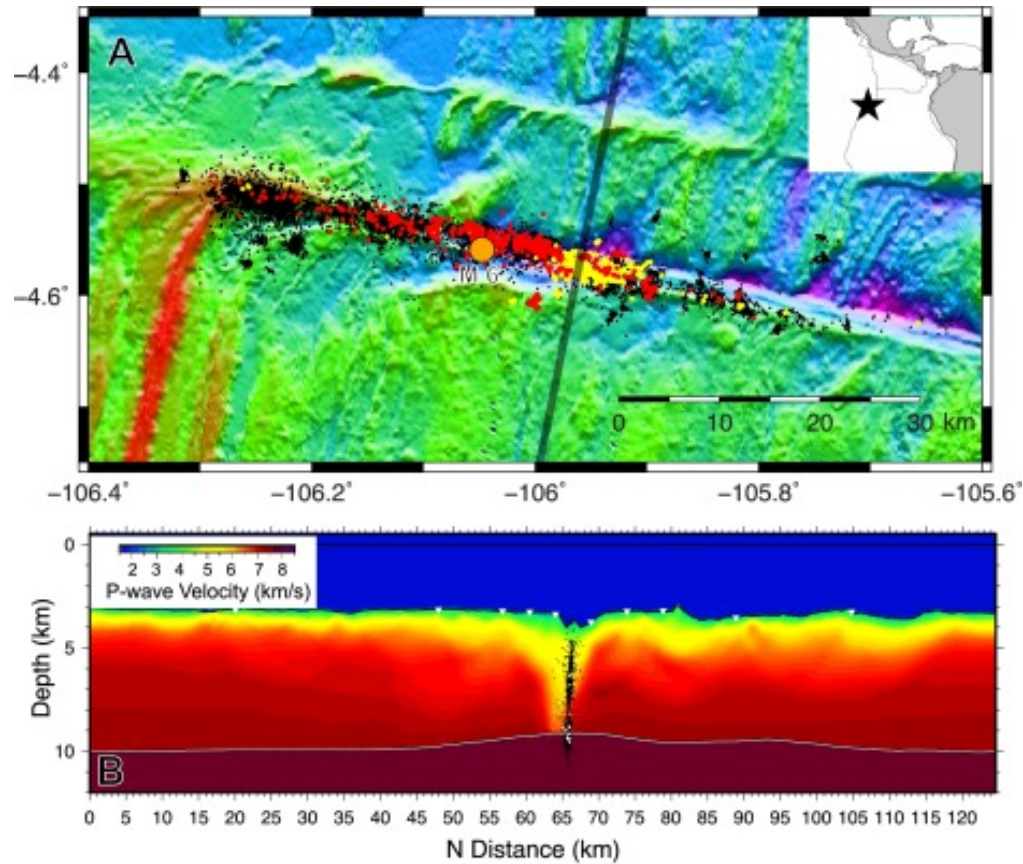


P wave and head wave travel times

Modeling of trapped waves



Fault zones are damaged

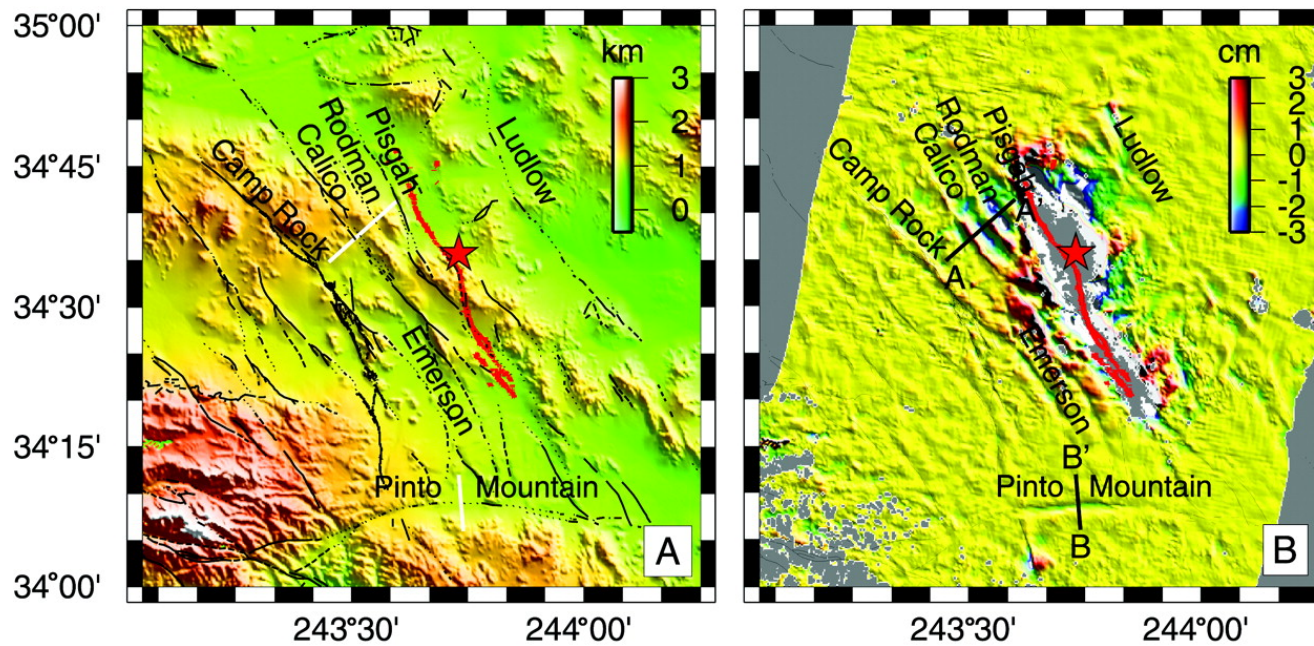


Central section of
the Gofar transform
fault on the East
Pacific Rise

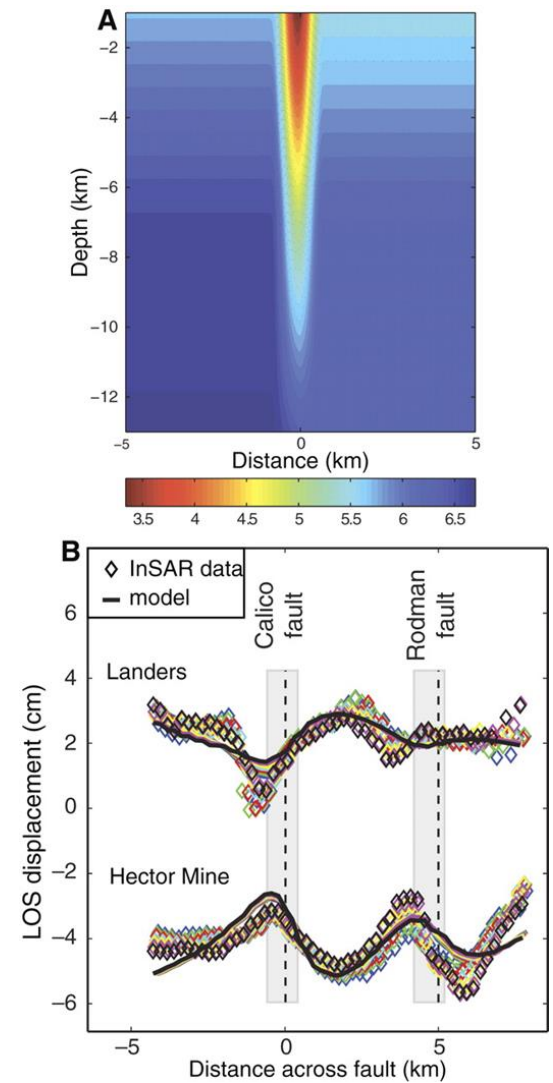
>2-km-wide low-
velocity fault zone
down to the base of
the crust

Roland et al (2012)

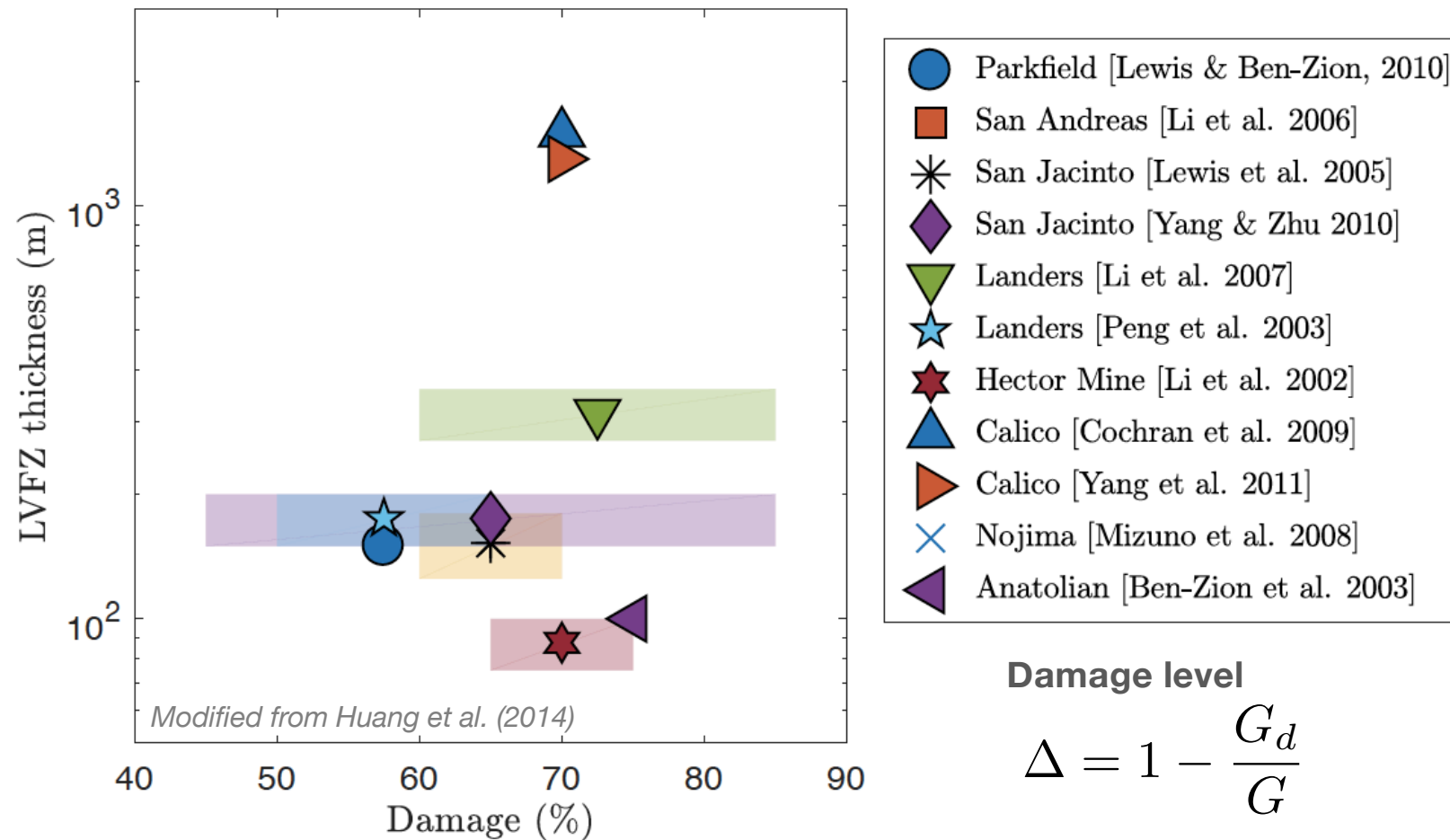
Fault zones are damaged



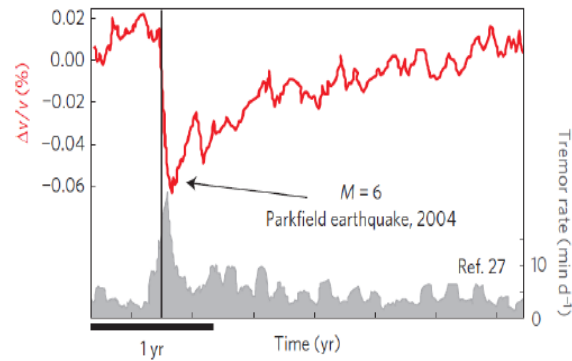
Fialko et al (2002), Cochran et al (2009)



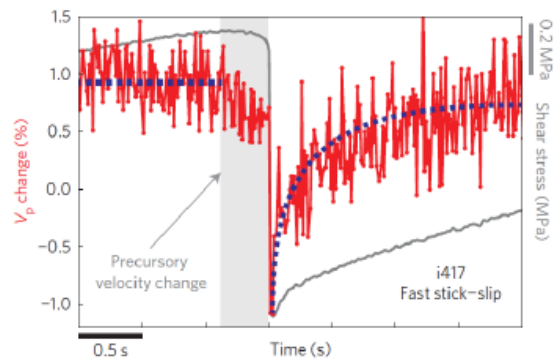
Fault zone properties



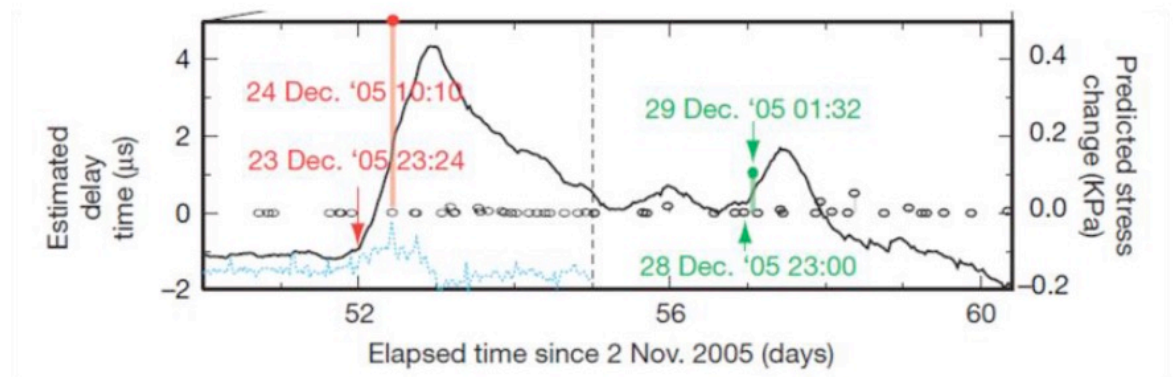
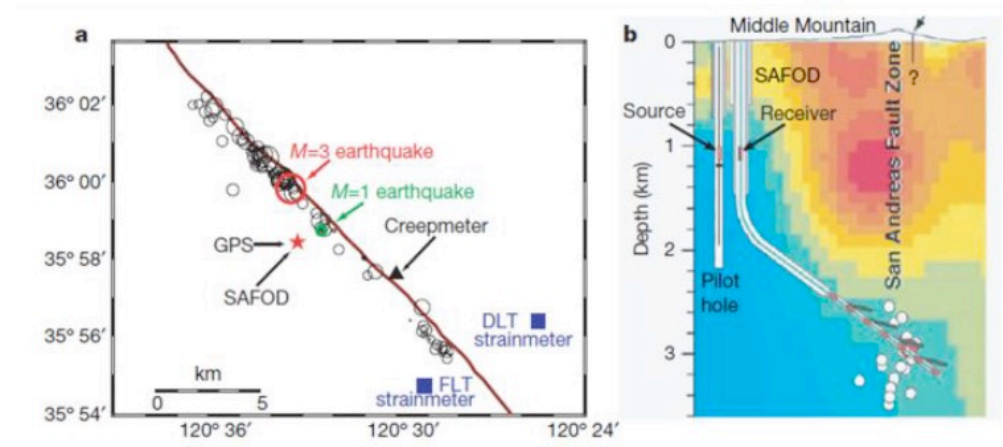
Crustal velocity changes post- and pre-seismic



Brenguier et al (2008)

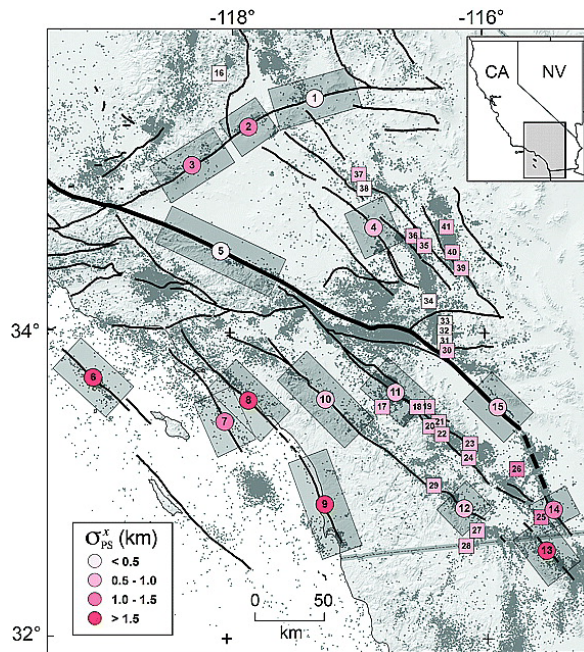


Scuderi et al (2016)

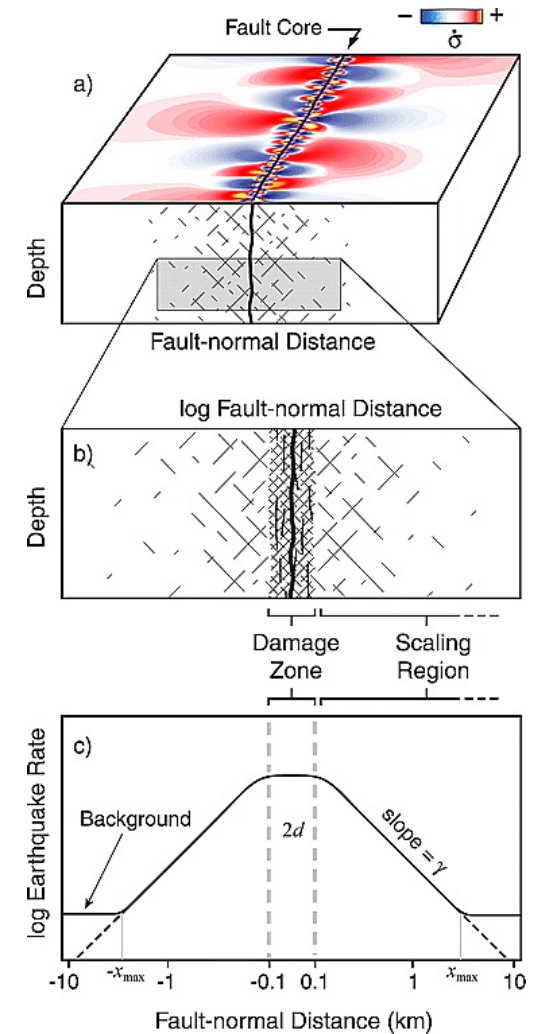
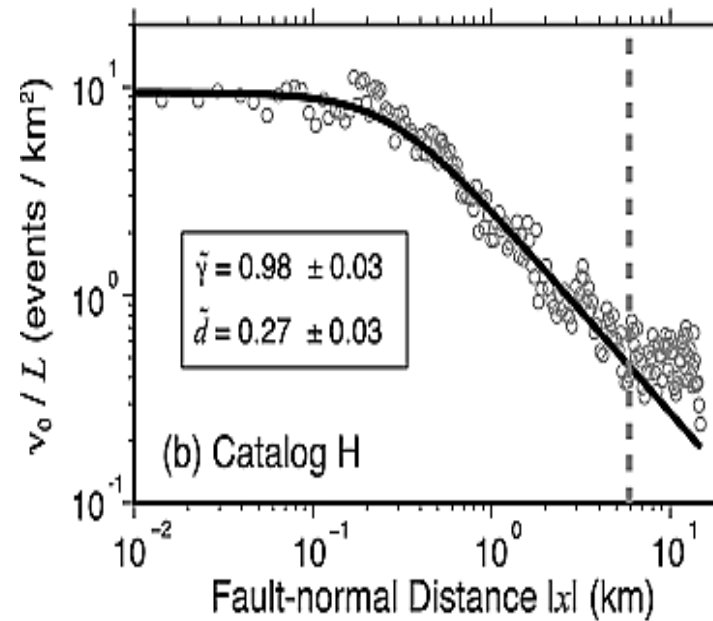


Niu et al (2008)

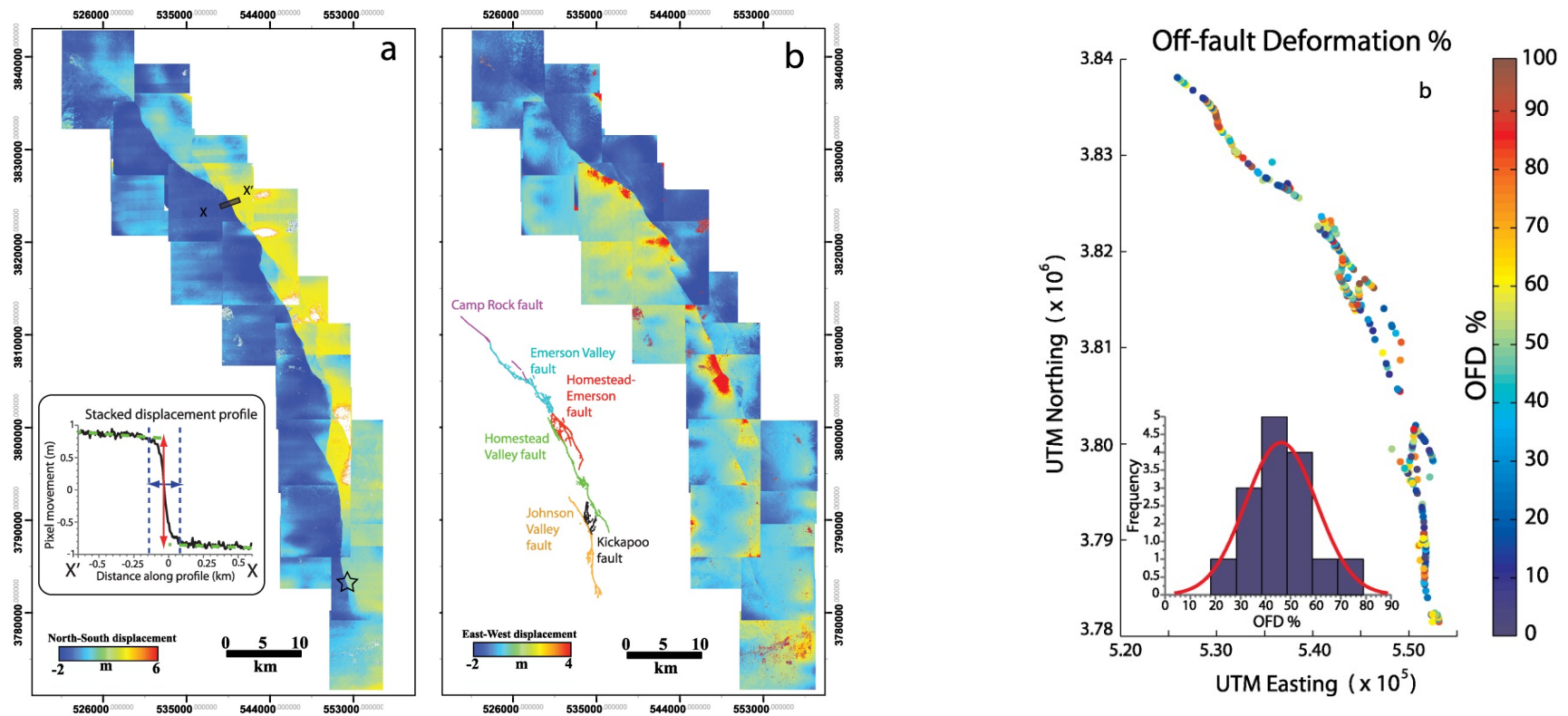
Fault zones are damaged



Distribution of seismicity in the vicinity of main faults in California
Powers and Jordan (2010)



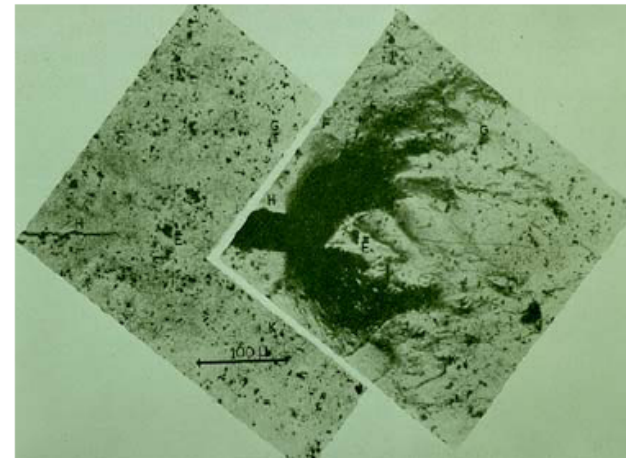
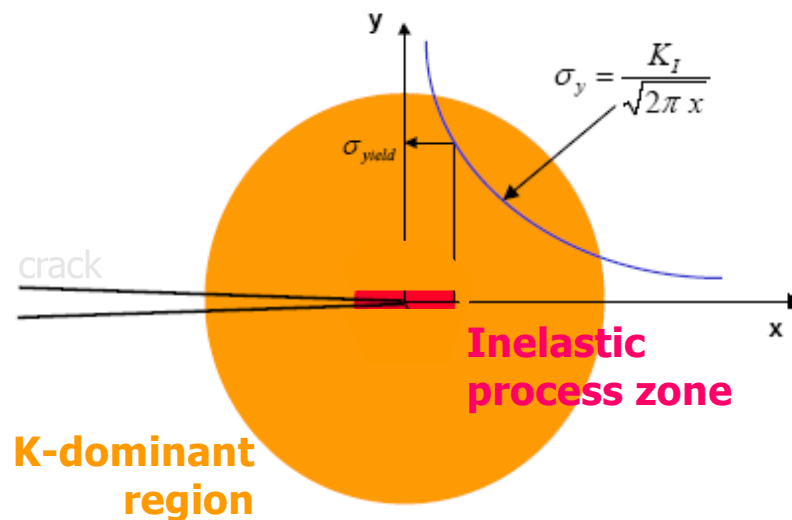
Coseismic damage



Optical satellite images, Landers earthquake, Milliner et al (2015)

A perspective from fracture mechanics

Linear elastic fracture mechanics (LEFM) predicts a stress singularity at the rupture front

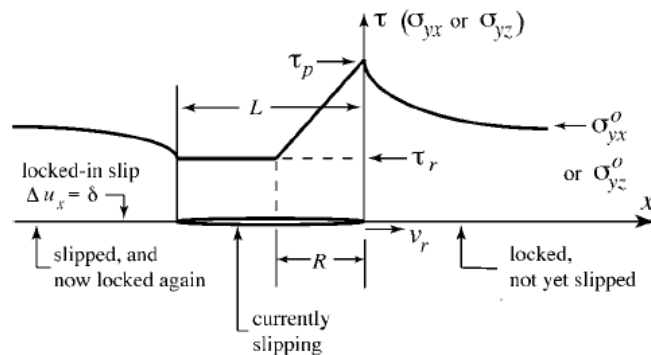


Process zone formation in Al-Cu-Mg alloy, From Broek, Elementary Engineering Fracture Mechanics, 3d Edit, 1982.

The stress concentration must be physically accommodated by nonlinear material behavior (damage, plasticity, micro-fractures)

Kostrov, Freund, Hussein, Kikuchi, Ida, Andrews (60-70s)

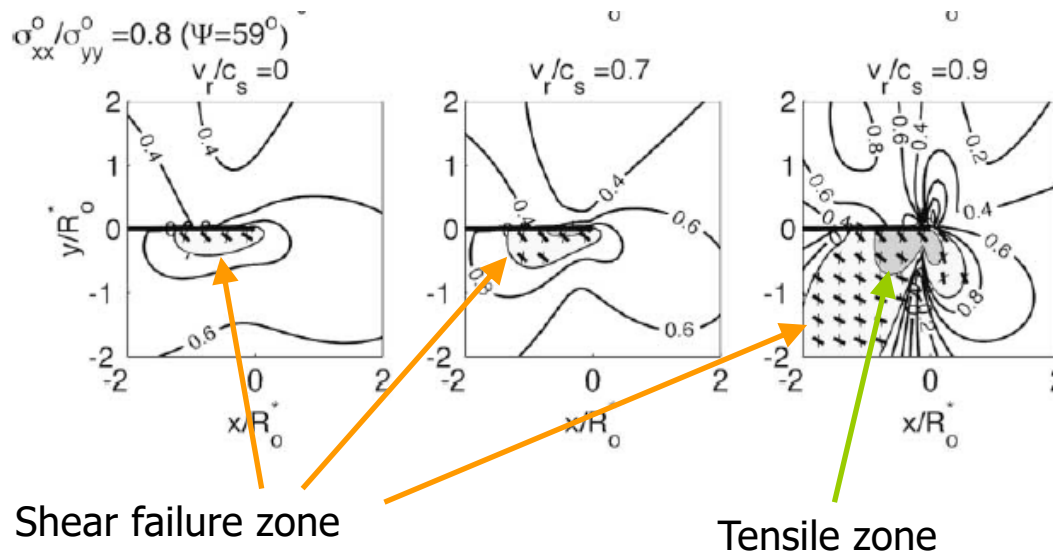
Predictions from a steady pulse rupture model
(Rice, Sammis and Parsons, 2005)



Characteristic size of off-fault damage zone:

$$R_o^* = \frac{9\pi F(0)}{16} \frac{\mu G}{(\tau_p - \tau_r)^2} = \frac{9\pi}{16(1 - \nu)} \frac{\mu G}{(\tau_p - \tau_r)^2}$$

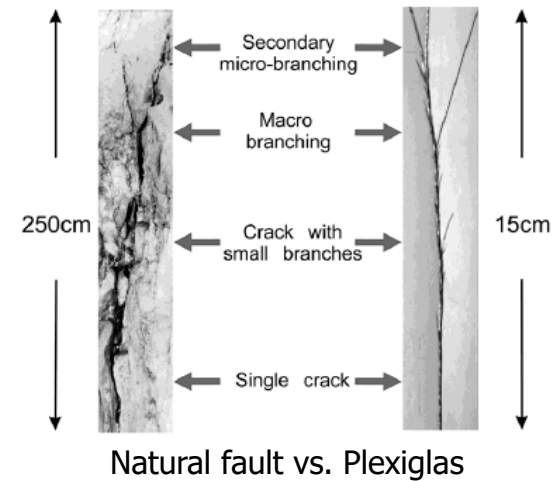
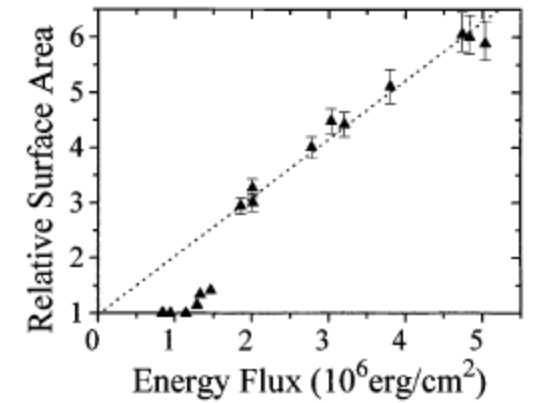
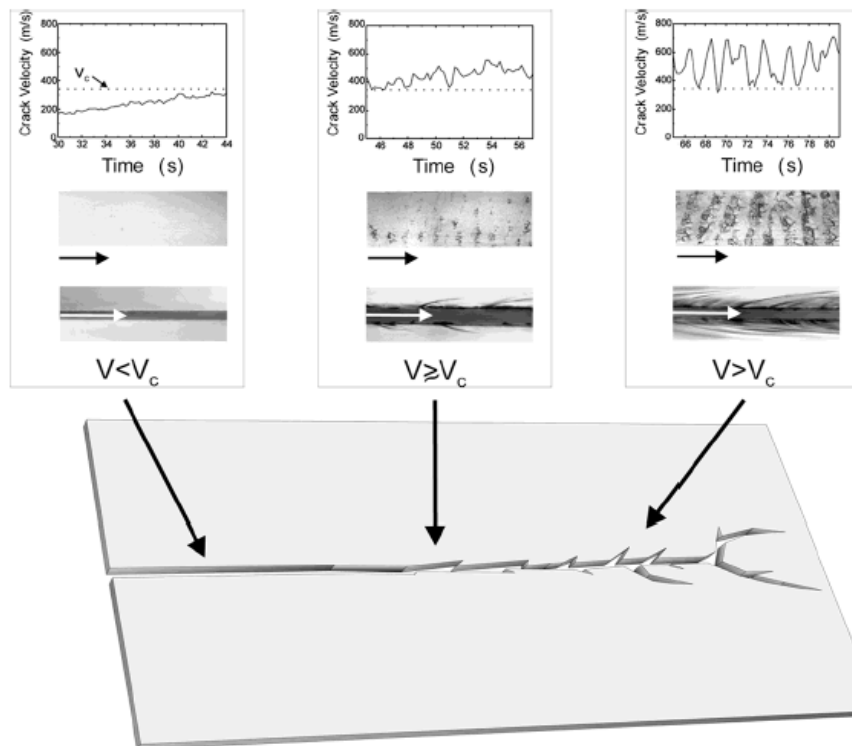
Estimates range from 1m to 1km



Contours of Coulomb stress outside the fault plane (>1 means failure)

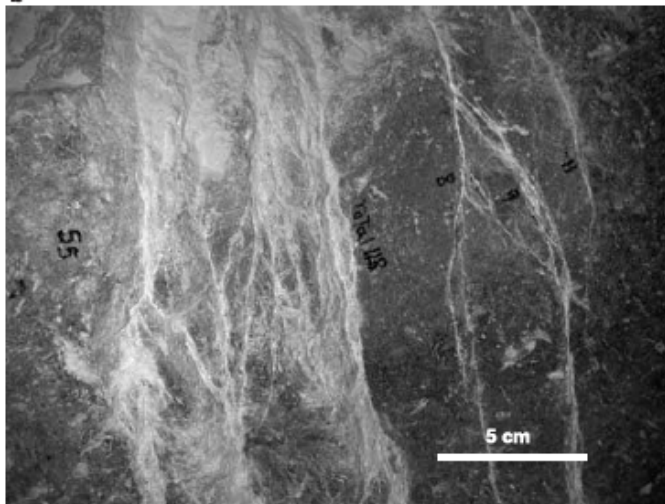
- Sharon, Gross and Fineberg (PRL, 1996) “Energy dissipation in dynamic fracture”
- Sharon and Fineberg (1999) “The dynamics of fast fracture”

Laboratory mode I rupture (opening) in plexiglas with controlled energy flux



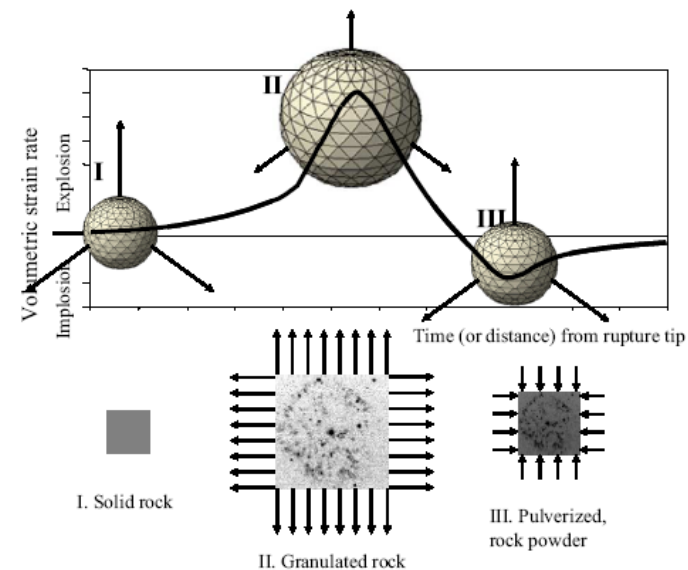
- Wilson et al. (Nature, 2005) “Particle size and energetics of gouge from earthquake rupture zones”
- Reches and Dewers (EPSL, 2005) “Gouge formation by dynamic pulverization during earthquake rupture”

Fracture zone of the Bosman fault, a **new** fault in a deep South African mine (M3.7, max slip 0.4 m): coalescing fractures filled with gouge powder



Measured grain size distribution
+multiple fracture branches

→surface energy $\sim 2\text{-}10 \text{ MJ/m}^2$



Modeling of secondary micro-cracks generated by dynamic ruptures

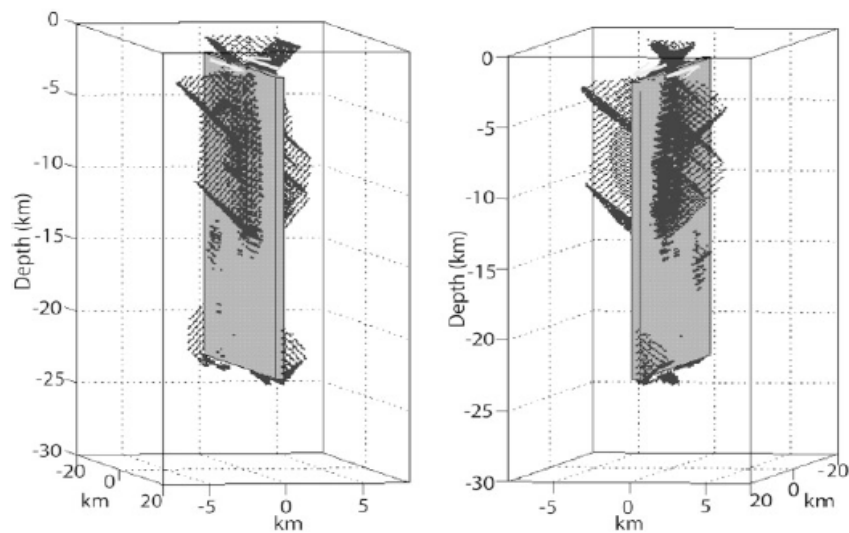
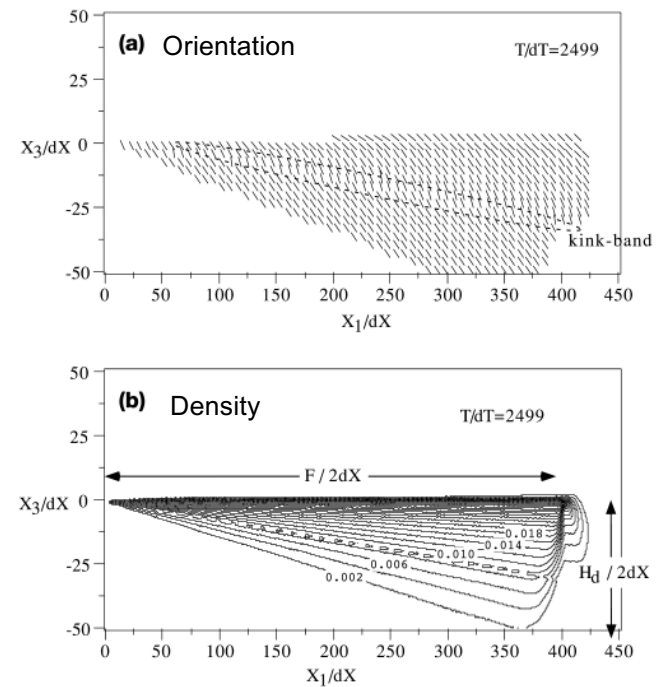


Figure 8. Two perspectives of the final stage of crack evolution for the 2000 Tottori earthquake dynamic simulation. The gray surface represents the shear crack on the pre-existing fault, and the dashed lines represents the tensile cracks.

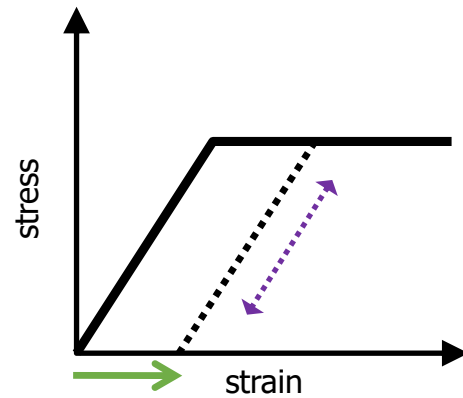
Dalguer et al (2003)



Yamashita (2000)

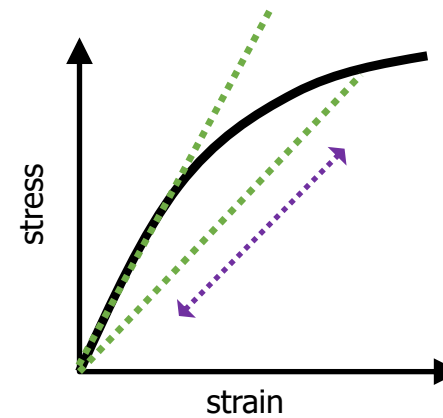
Continuum damage vs. plasticity

Plasticity



Plasticity describes the
generation of irreversible
strains

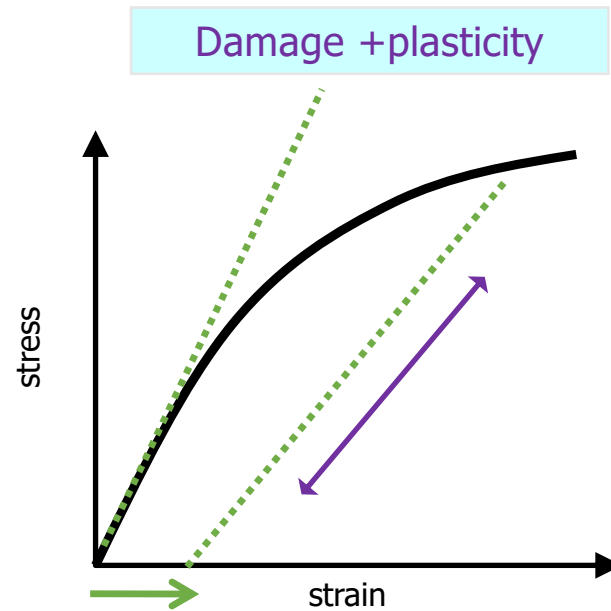
Damage



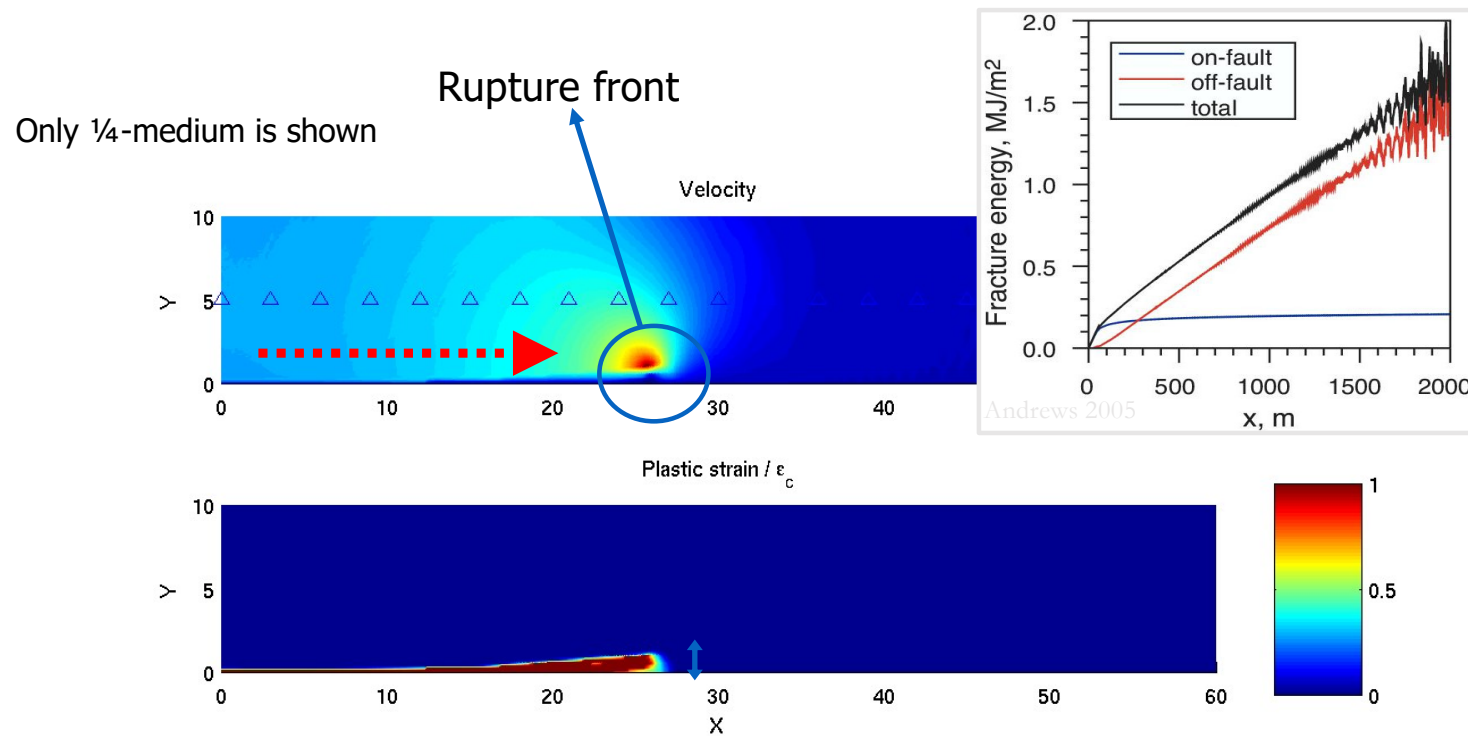
Damage describes changes in elastic
moduli due to microcracking

Continuum damage and plasticity

Both can be combined:
irreversible strain and reduction of elastic moduli



Strain weakening visco-plasticity outside the fault plane



Andrews
2005

The thickness of the dissipation zone increases as the rupture grows
→ the “apparent” fracture energy increases with rupture length

Stable rupture speed and process zone size

Crack tip equation of motion:

$$G_c = g(v)G_0 = g(v) \pi a \frac{\Delta \tau^2}{2\mu}$$

Off-fault dissipation \rightarrow scale-dependent total fracture energy $G_c = \gamma a$

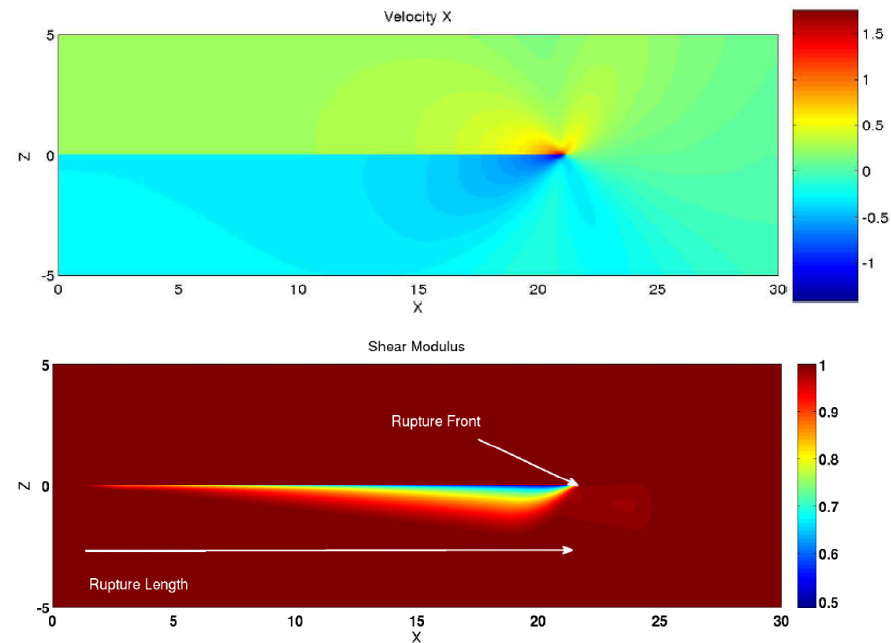
\rightarrow Steady rupture speed: $v = g^{-1} \left(\frac{2\gamma\mu}{\pi\Delta\tau^2} \right)$

\rightarrow Steady process zone size: $\Lambda = \sqrt{1 - v^2/\beta^2} \Lambda_0$

Continuum damage outside the fault

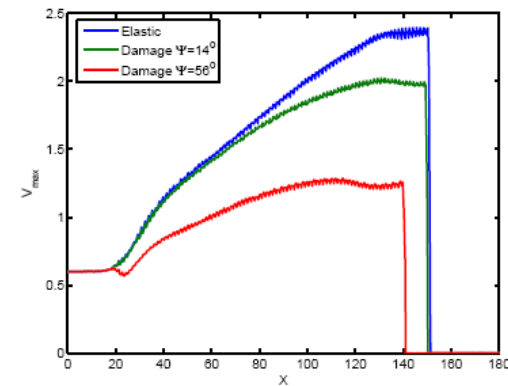
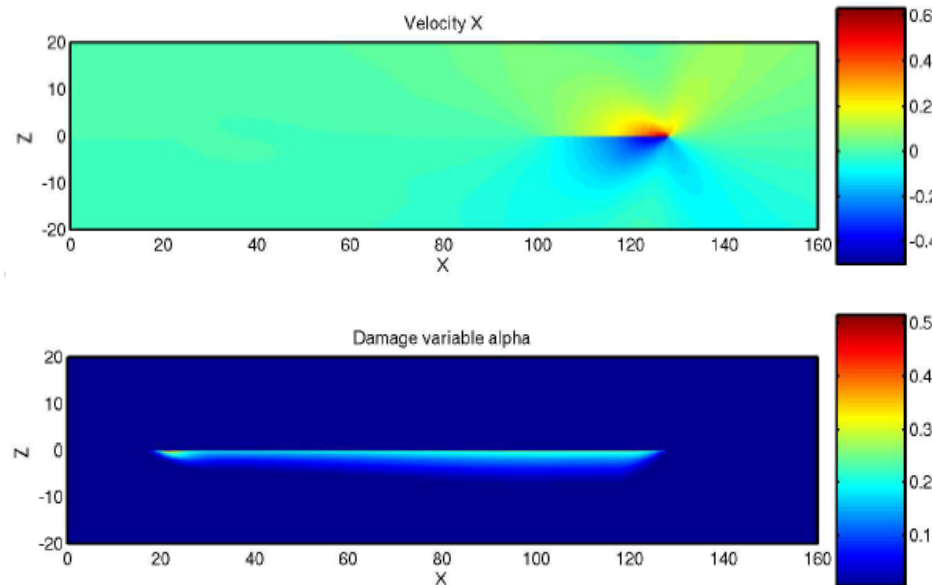
Dynamic simulations with a spectral element code (Ampuero et al, 2008)

Self-similar crack ruptures develop growing damage zones



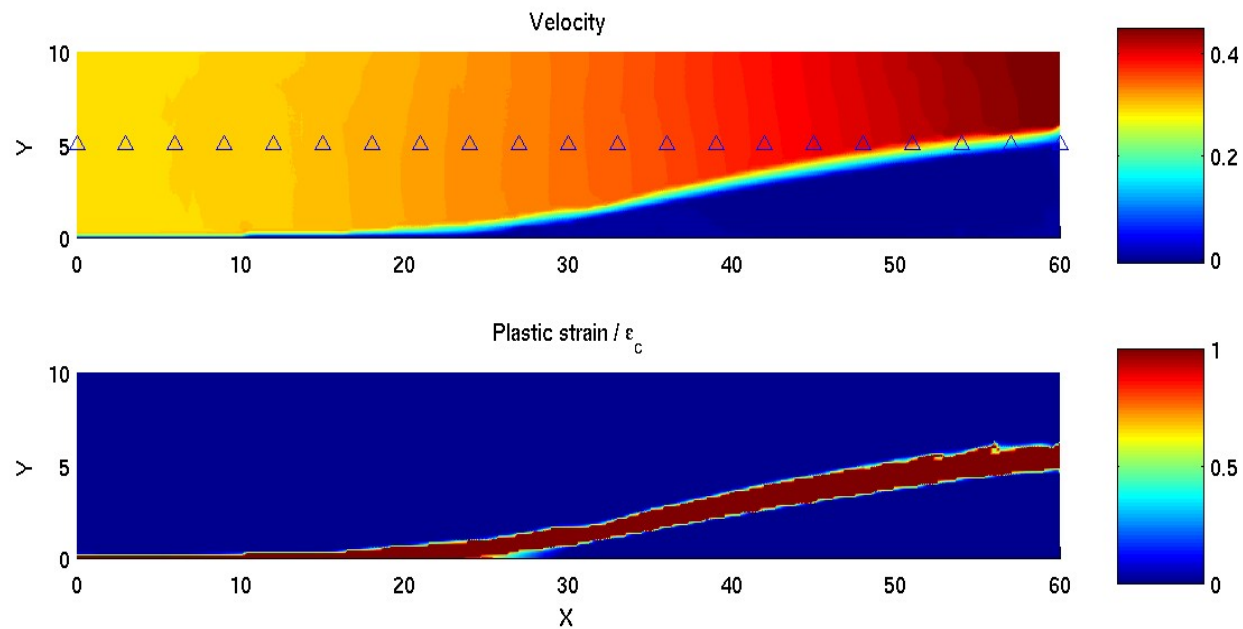
Continuum damage outside the fault

Self-similar pulse-like ruptures show similar features, but thinner damage zones controlled by the pulse width (rise time)



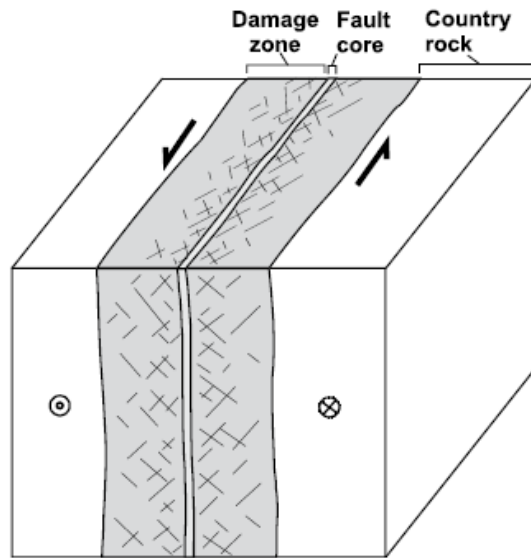
Dynamic off-fault damage reduces peak ground motions (Ampuero et al, 2008)

Strain weakening visco-plasticity outside the fault plane



Interesting problem: rupture branches out spontaneously
when not guided by a weak fault plane

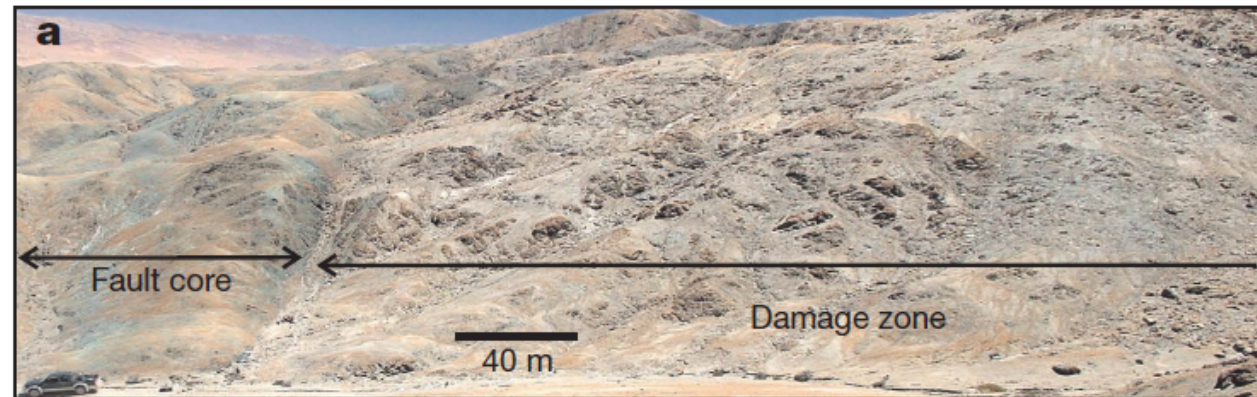
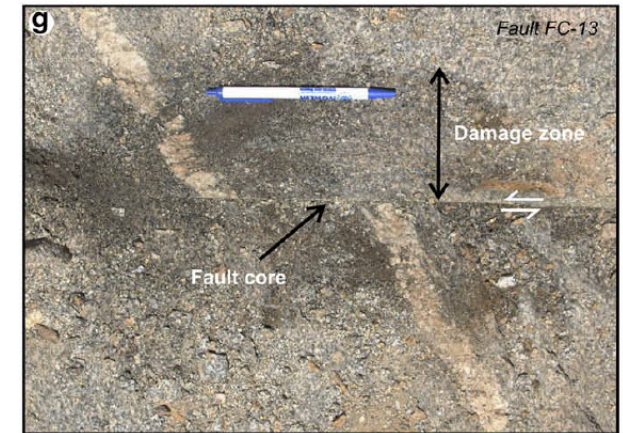
How do ruptures generate
fault damage zones?



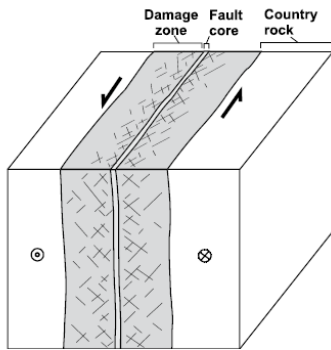
fault zone

Mitchell and Faulkner (2009)

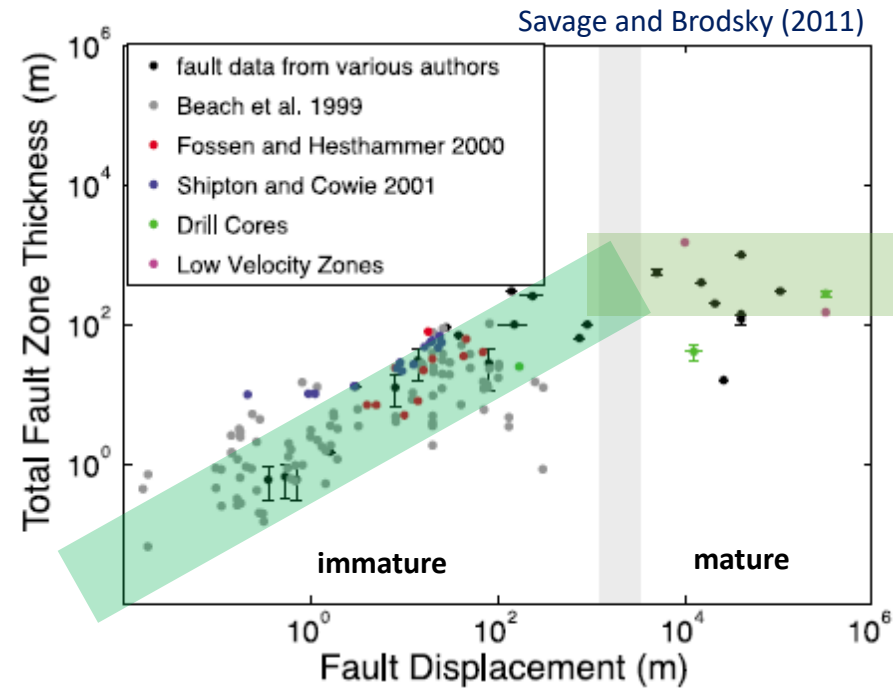
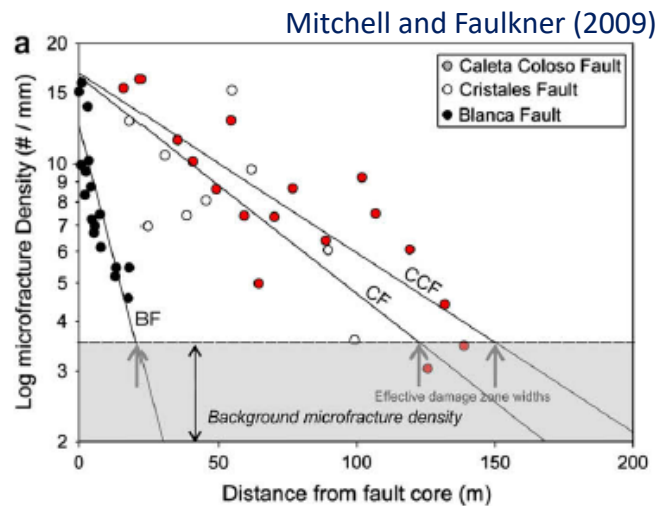
Damage zone thickness varies from a few centimeters to several hundred meters.



Faulkner et al. (2006)



Damage zone thickness saturates
at large fault displacement

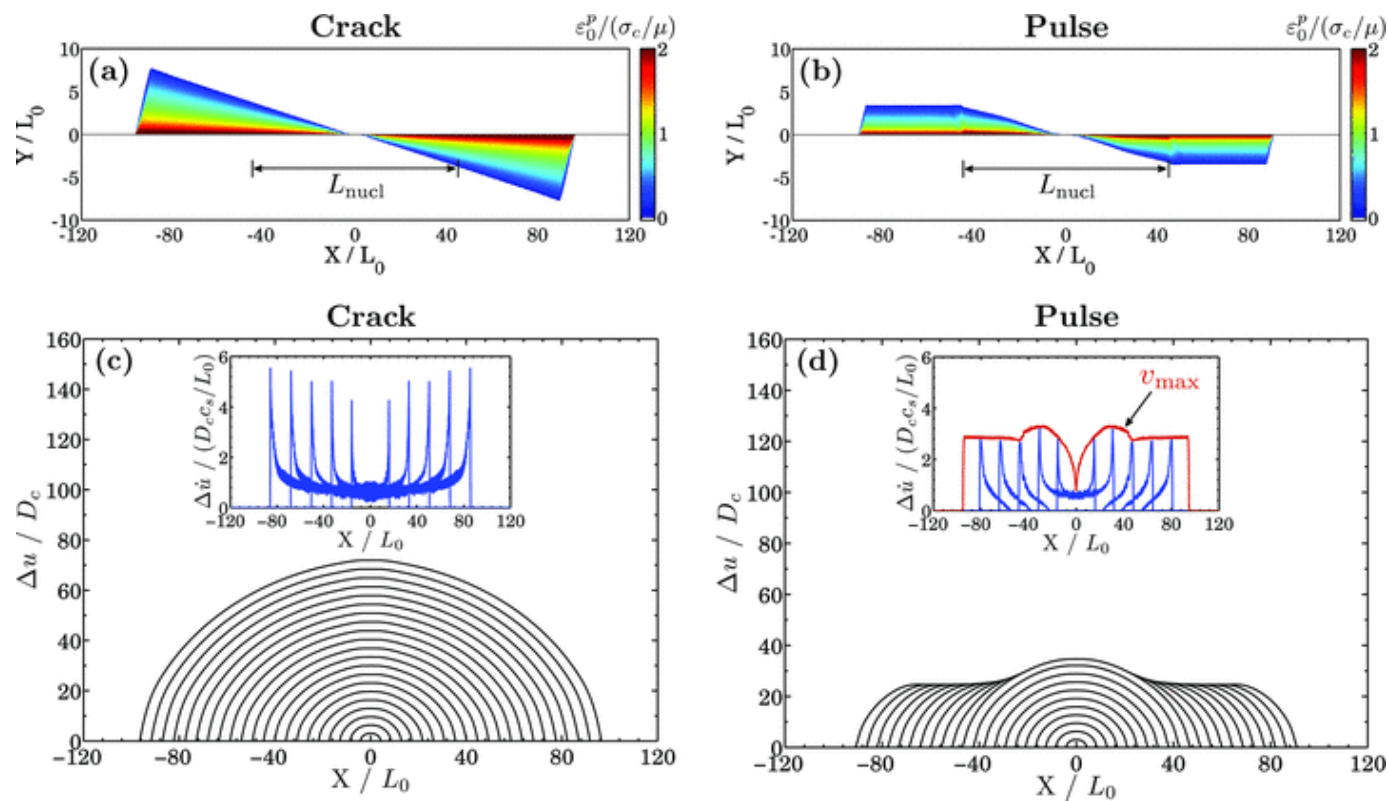


$< 1 \text{ km}$

What limits the thickness of damage zones ?

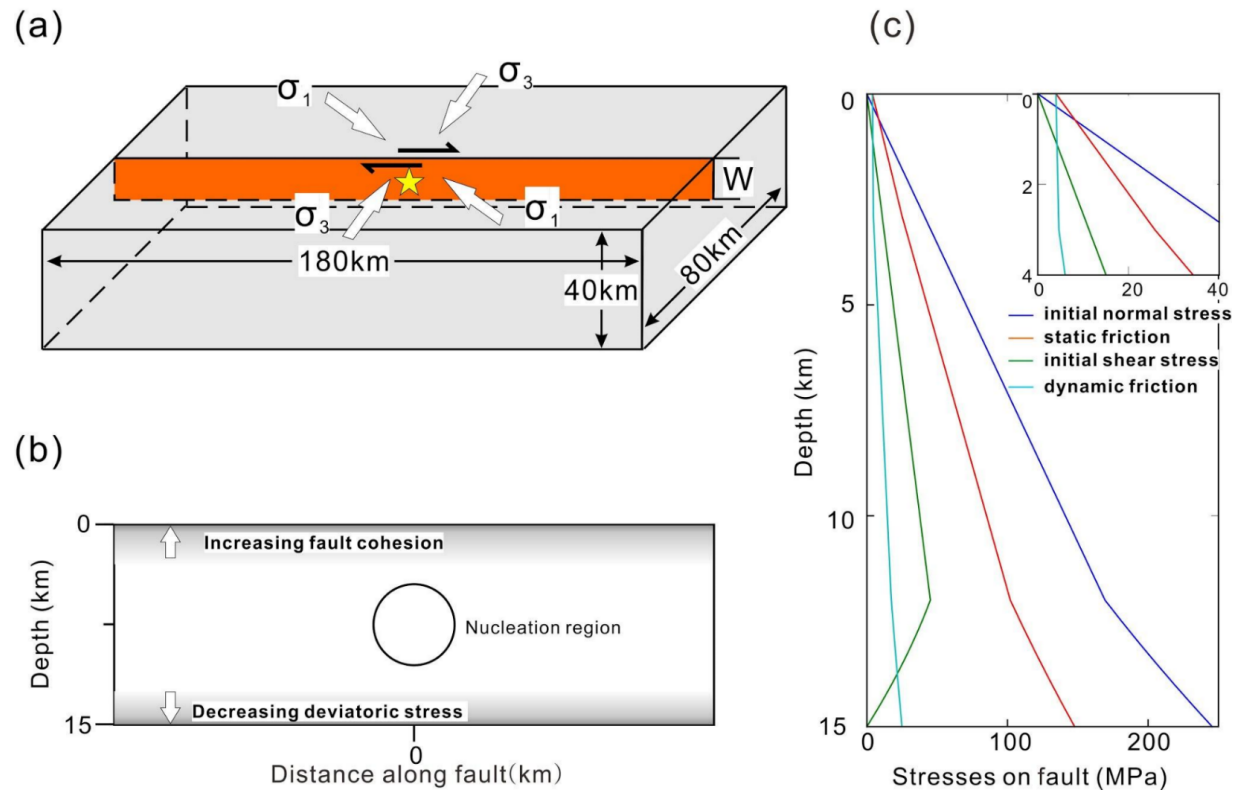
Ampuero and Mao (2017), Upper Limit on Damage Zone Thickness Controlled by Seismogenic Depth

Coseismic off-fault inelastic deformation

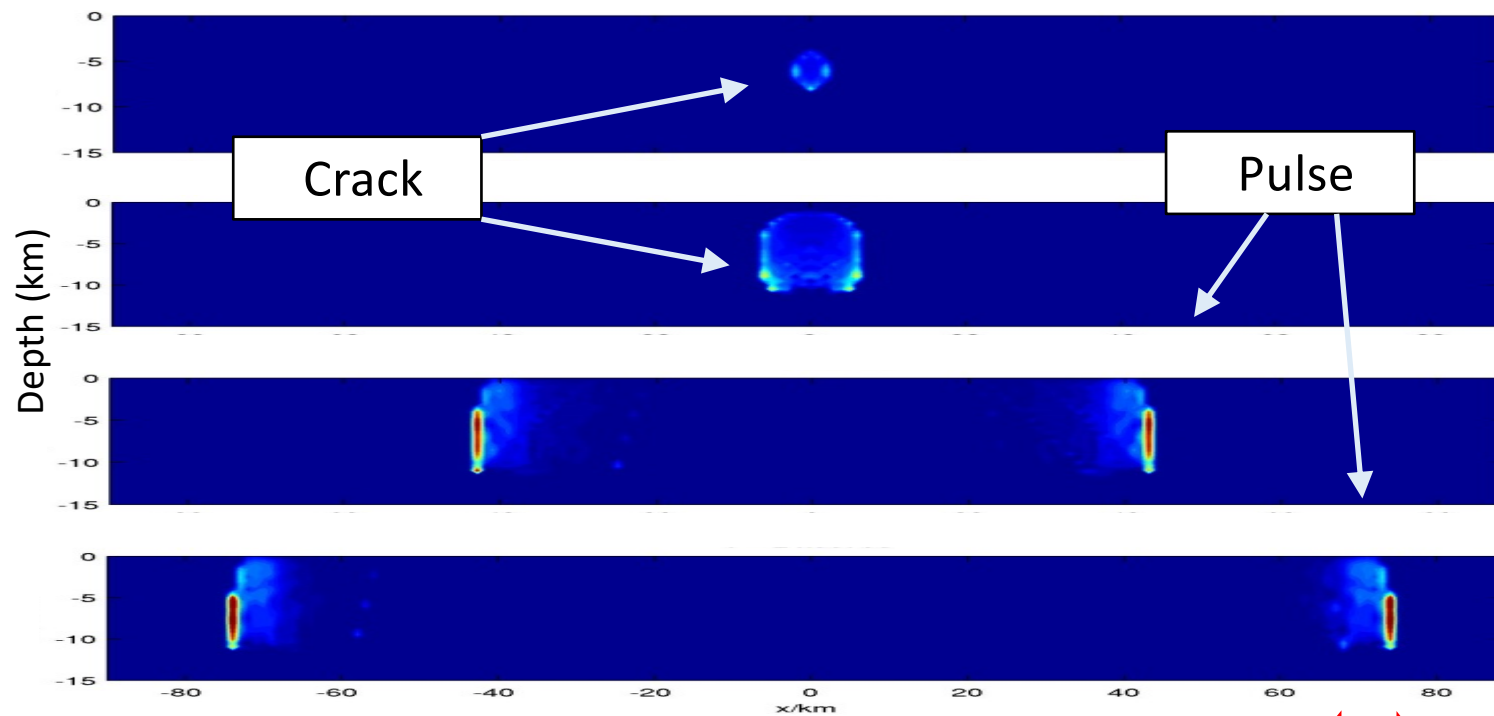
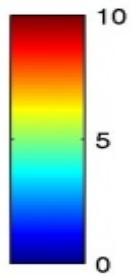


Xu, Ben-Zion and Ampuero (2014)

Dynamic rupture models with coseismic off-fault damage



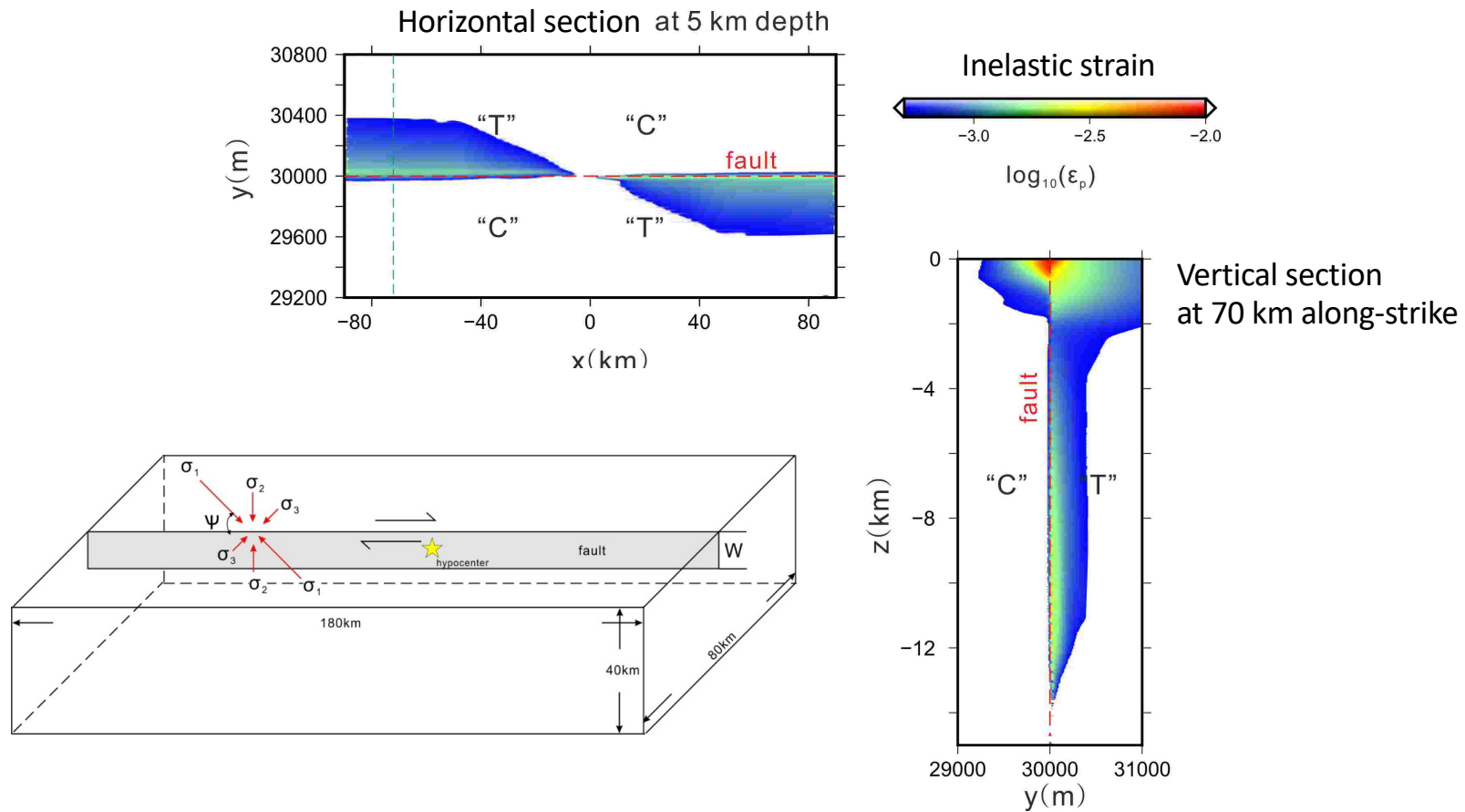
Slip rate
(m/s)



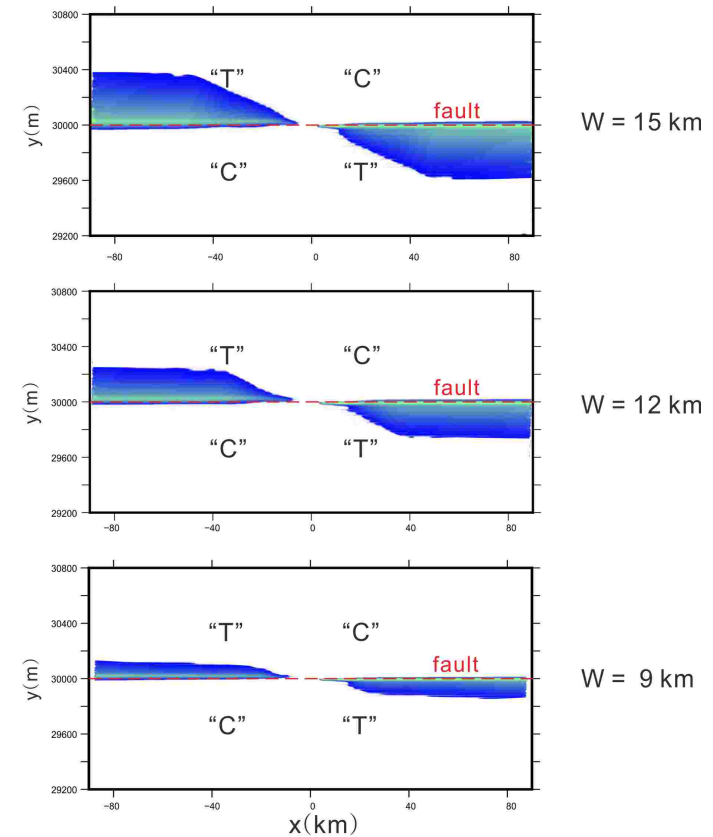
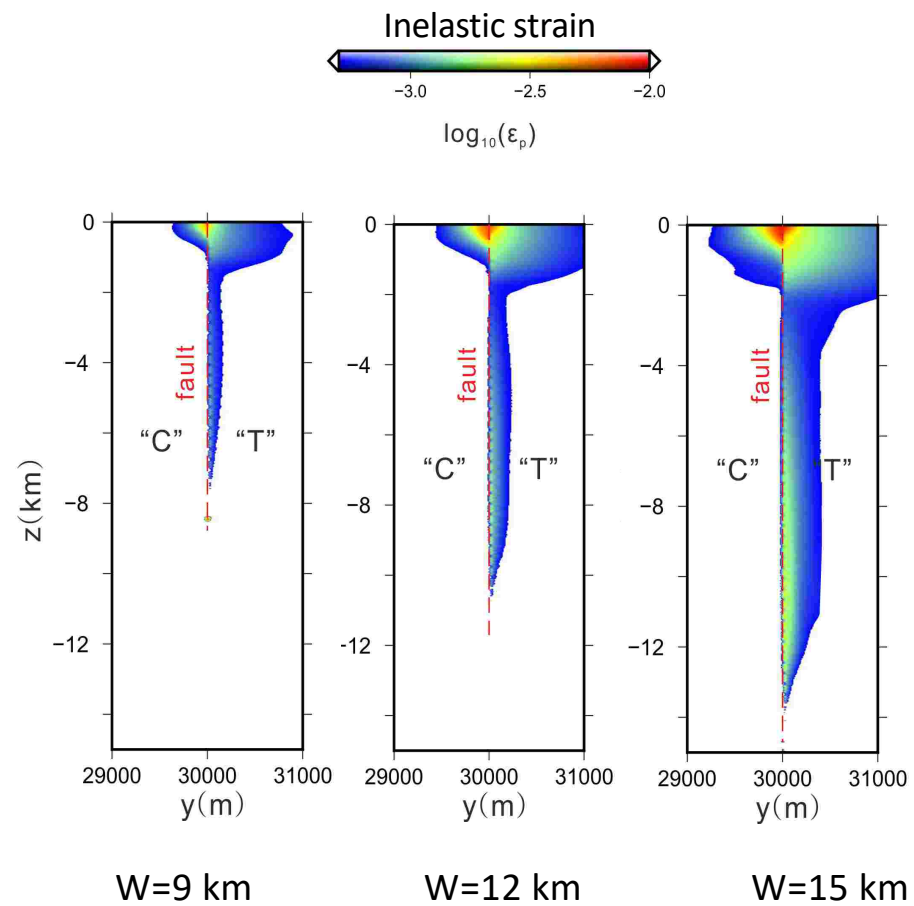
Seismogenic
thickness

Pulse width

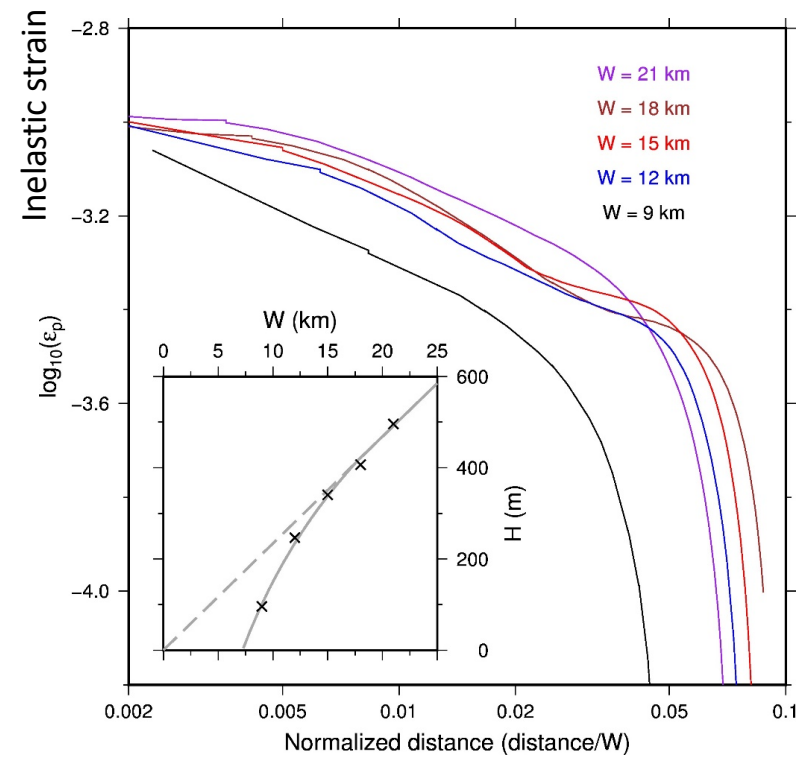
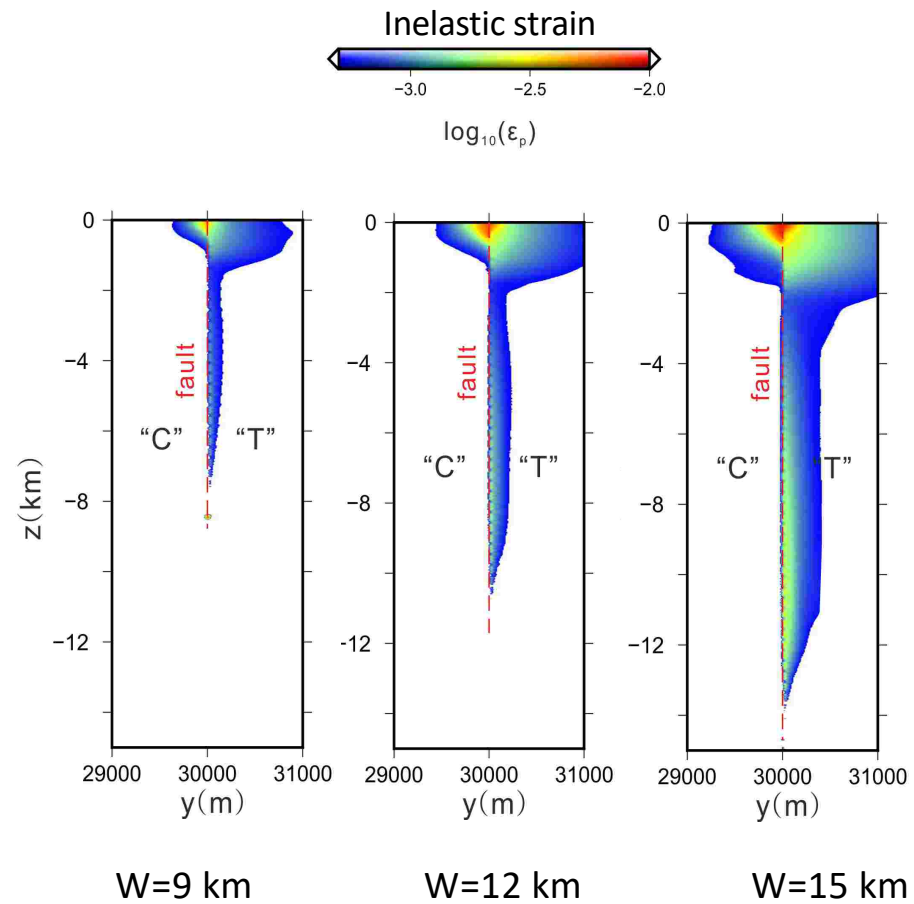
Plastic strain distribution (W=15 km)



Inner damage zone thickness depends on seismogenic width

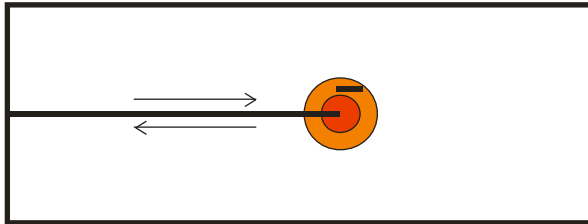


Inner damage zone thickness depends on seismogenic width

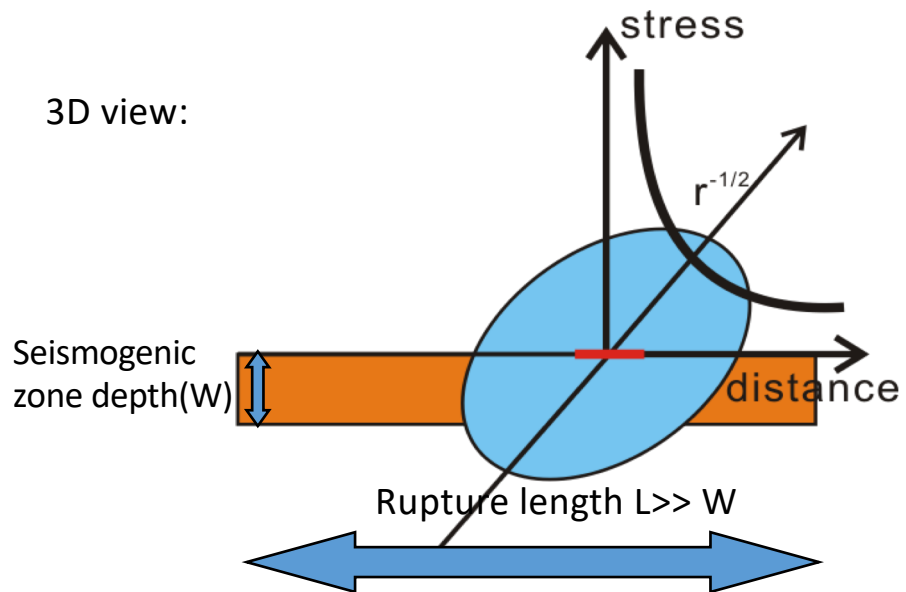


Fracture mechanics theory

Map view:



3D view:



Stress near crack tip: $\tau \approx \frac{K}{\sqrt{r}} + \tau_0$

where K is the stress intensity factor, $K \sim \sqrt{l} \Delta\tau$

$\Delta\tau$ is stress drop and l the shortest rupture size:

$l = R$ (radius) for circular ruptures,

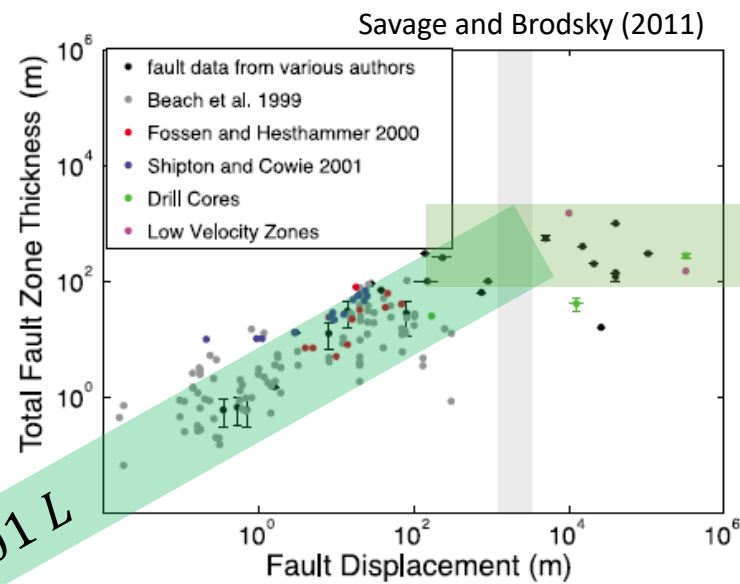
$l = W$ (width) for elongated ruptures ($W \ll L$)

Damage zone size: distance at which $\tau = \tau_s$
(stress=yield strength)

$$r_c \sim \left(\frac{\Delta\tau}{\tau_s - \tau_0} \right)^2 l < \sim 0.01 W$$

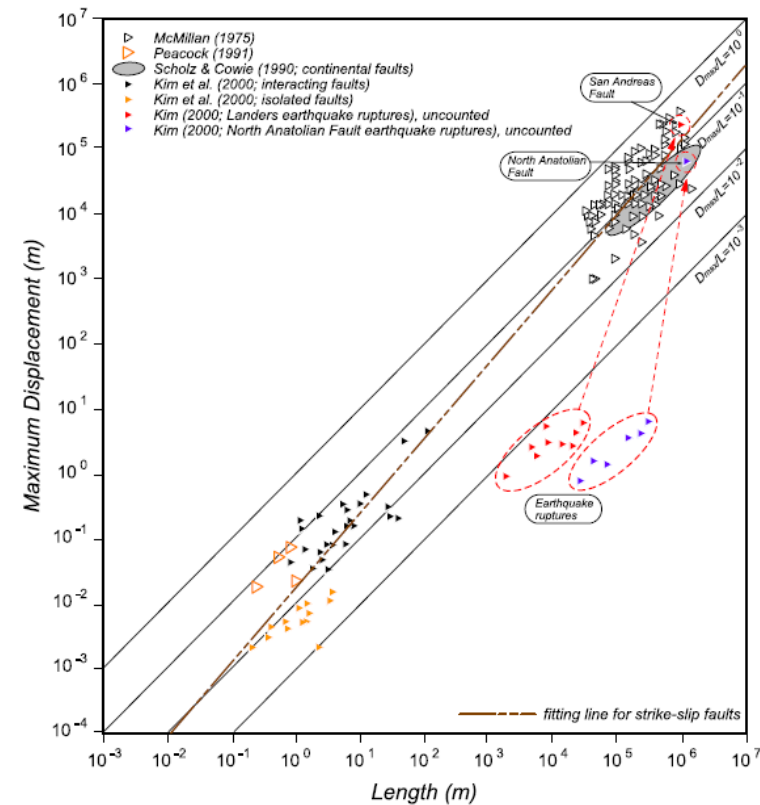
Relative stress drop

Inner damage zone thickness



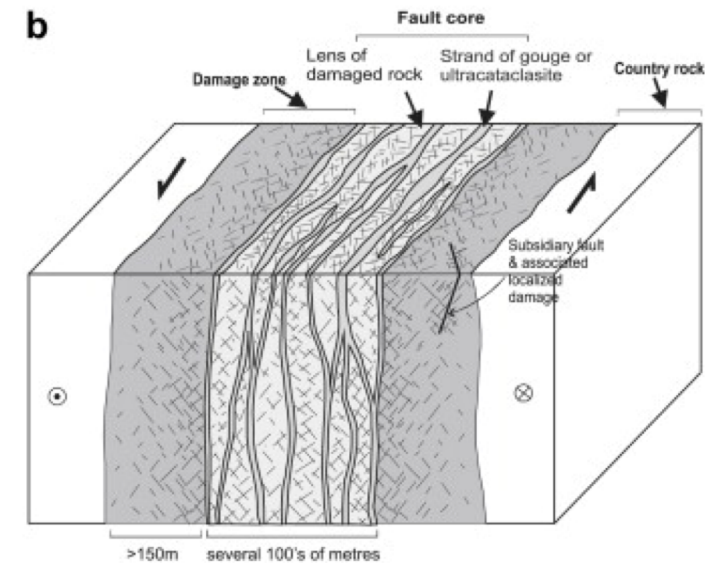
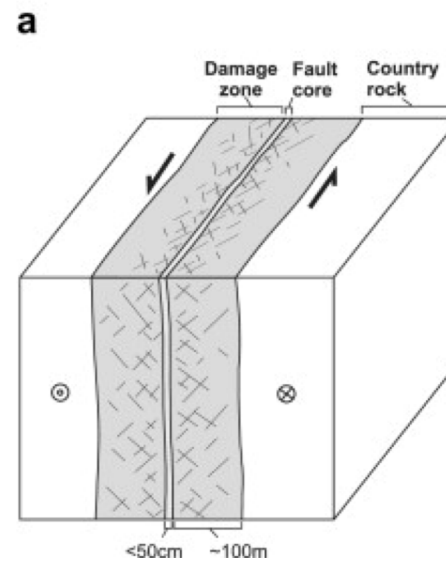
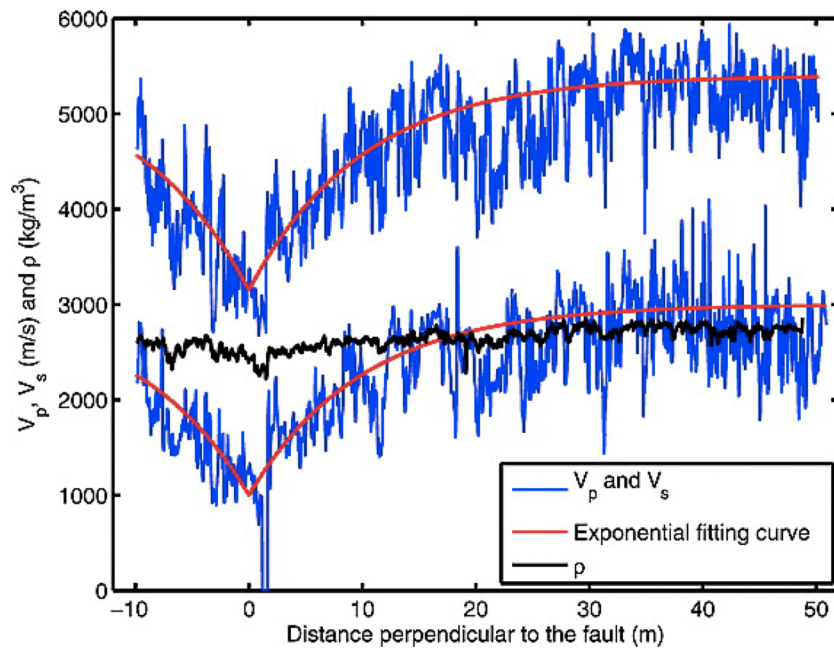
$$r_c \sim 0.01 L$$

$$\sim (0.01 - 0.1) \times \text{Length}$$



How does pre-existing damage
affect ruptures?

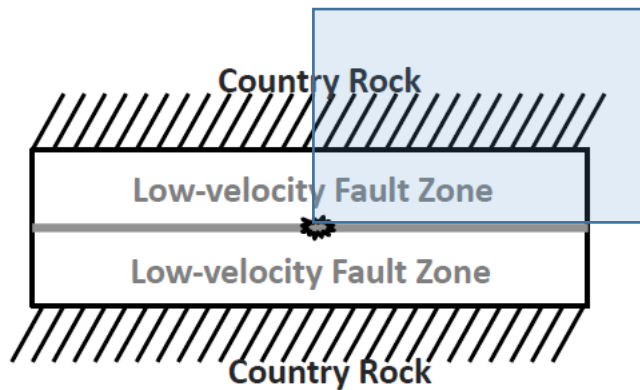
Low velocity fault zones in nature



Mitchell & Faulkner 2009

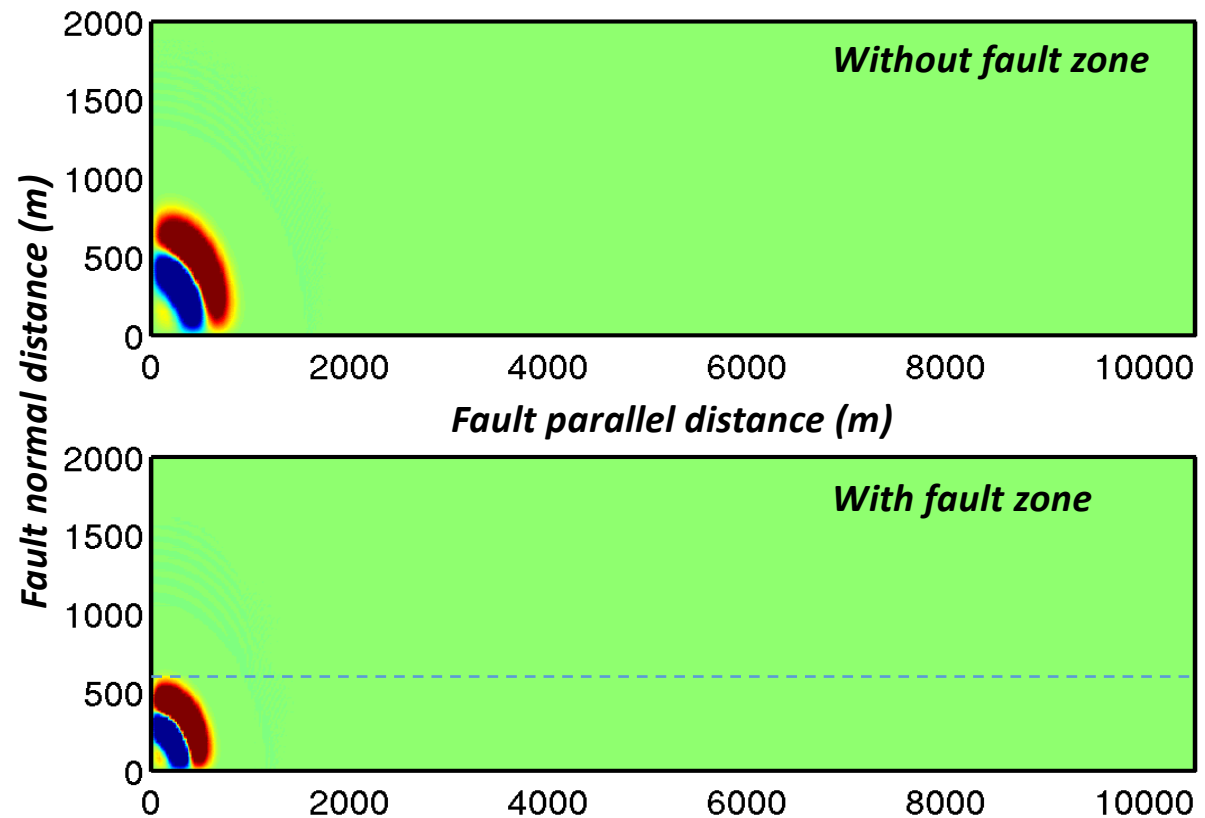
Huang & Ampuero 2011

Fault zone damage affects dynamic rupture

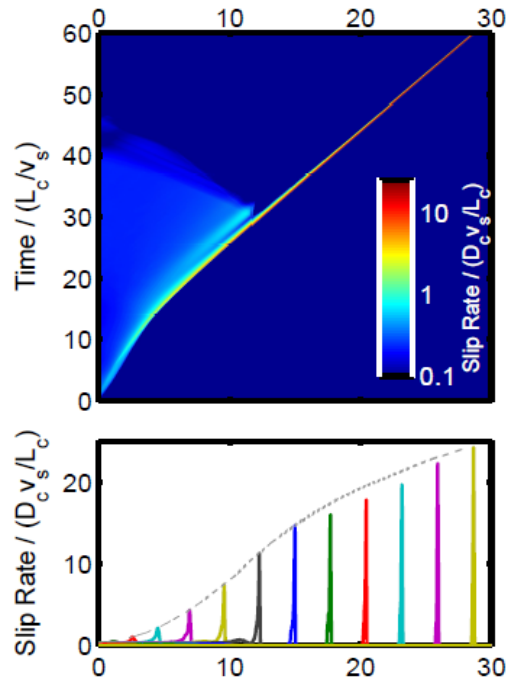


Dynamic rupture simulations on faults bisecting a damaged (low velocity) zone

Huang and Ampuero (2011)
Huang, Ampuero and Helmberger (2014, 2015)

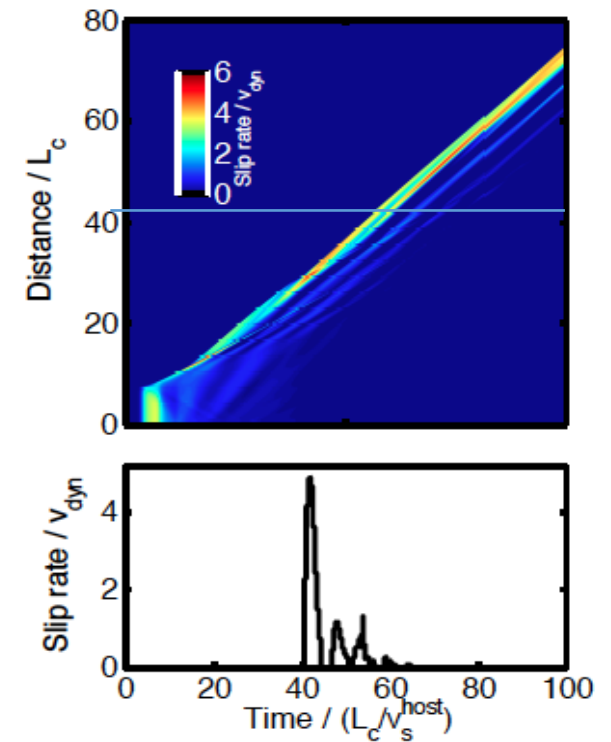


Fault zone damage affects dynamic rupture



Fault zone effects:

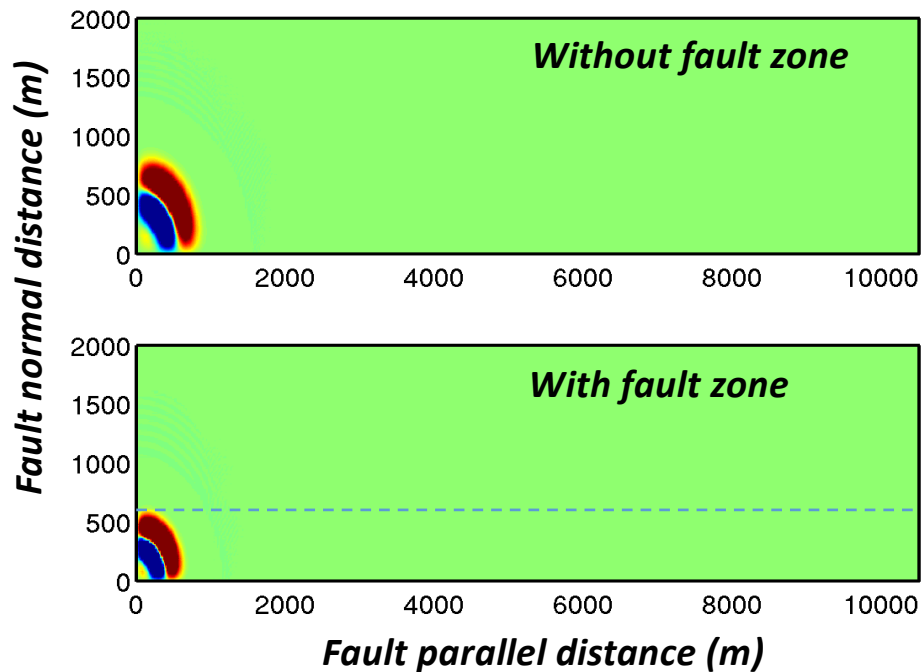
- Short rise time (slip pulses)
- Oscillatory slip rate (multiple pulses)
- Oscillatory rupture speed
- Supershear transition at low stress
- **Rupture at unexpected speeds**
- Spatially periodic off-fault damage patterns
- **Holes in radiated spectra**



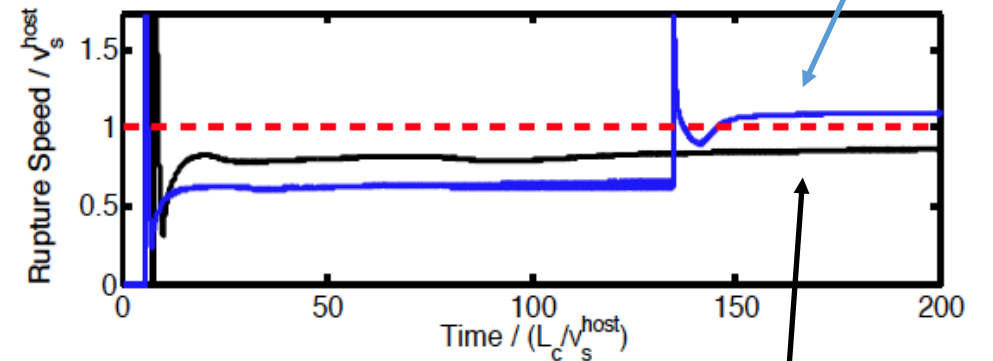
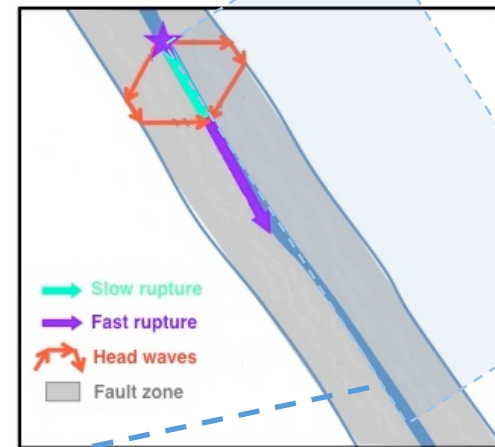
Huang and Ampuero (2011)

Huang, Ampuero and Helmberger (2014)

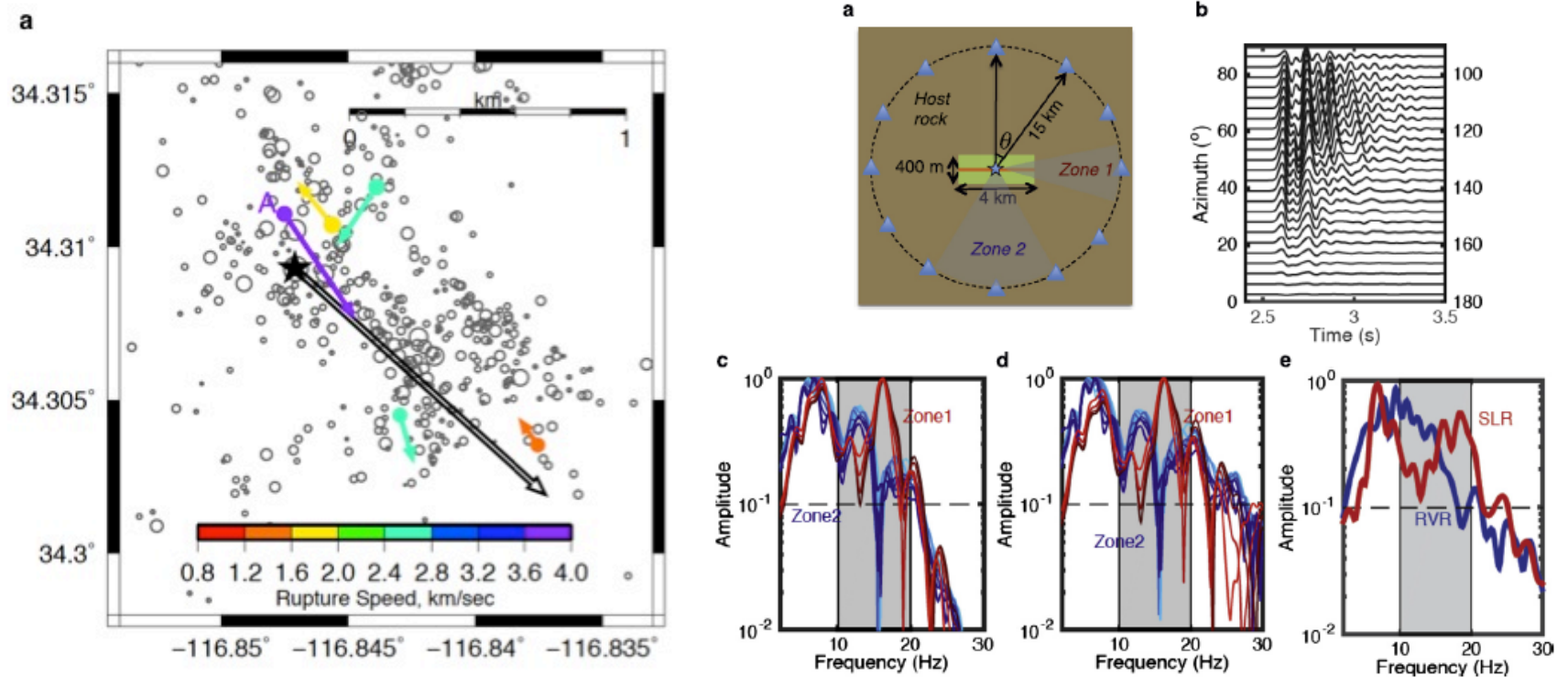
Fault zone waves facilitate supershear transition



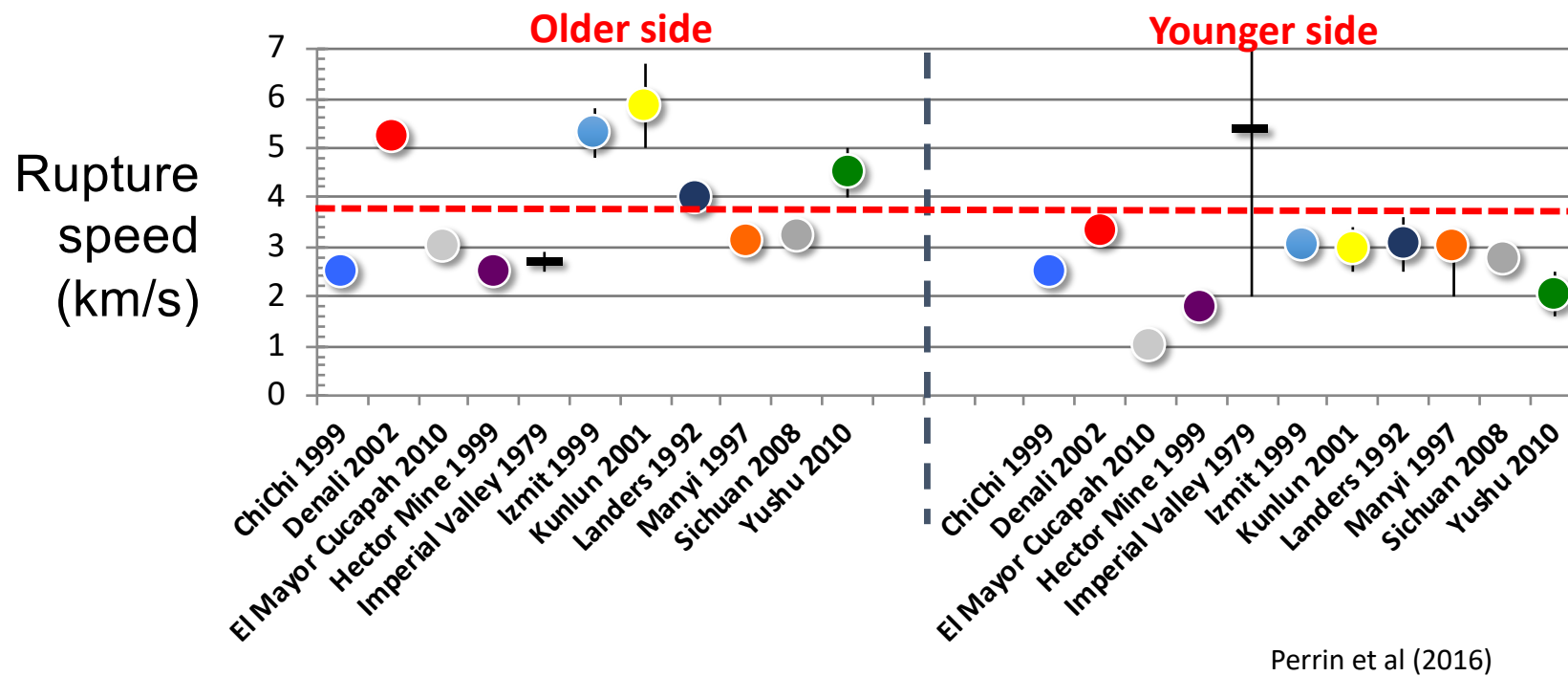
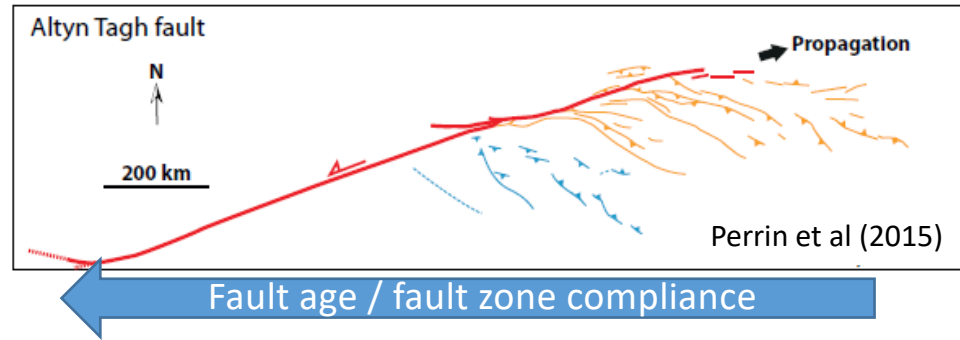
Huang and Ampuero (2011)
Huang, Ampuero and Helmberger (2014, 2015)



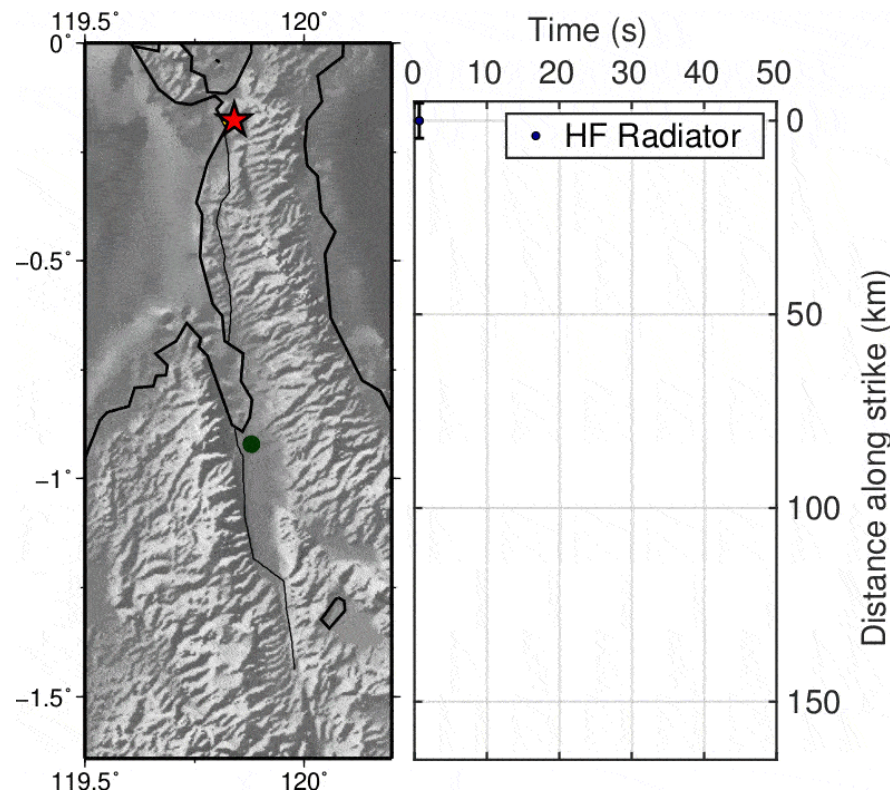
Evidence of rupture speed enhanced by fault zones in microearthquakes



Huang, Ampuero and Helmberger (2015)



Fast rupture of the 2018 Mw 7.5 Palu earthquake



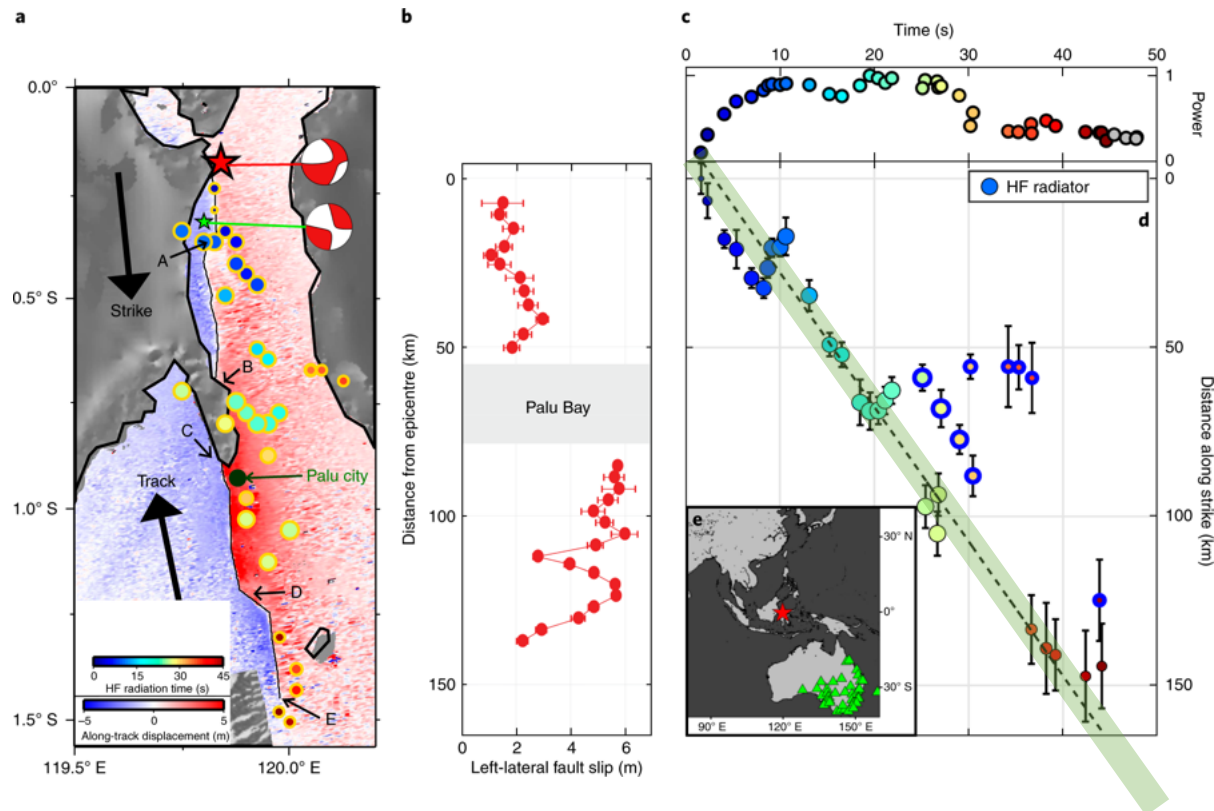
Teleseismic observations of rupture speed by back-projection rupture imaging

Southward rupture length ~150 km

A **supershear** earthquake:
Rupture speed faster than S waves

Steady and fast rupture despite large fault bends

Teleseismic observations of rupture speed



Rupture speed ≈ 4.1 km/s

Fast speed from early on

Steady despite fault bends

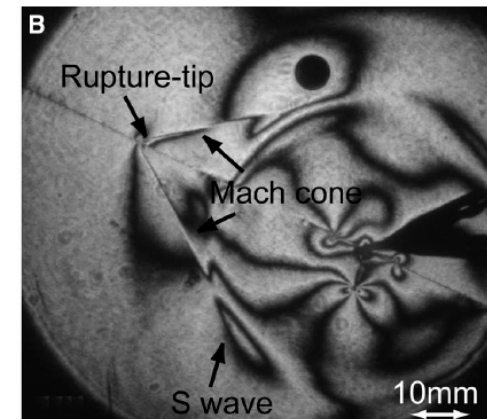
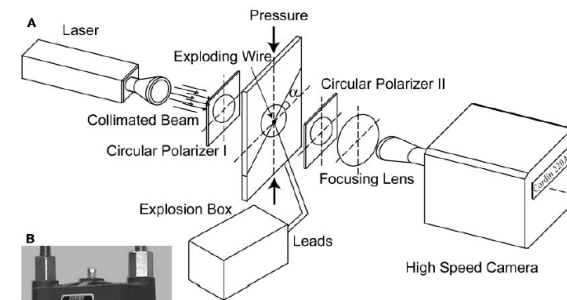
Teleseismic observations of rupture speed

2. Surface wave Mach cone



Sonic boom of a super-sonic jet plane

Supershear laboratory earthquake experiments in Ares Rosakis group
(Caltech) (Xia *et al*, 2004)



Teleseismic observations of rupture speed

2. Surface wave Mach cone



Sonic boom of a super-sonic jet plane
(body wave analogy)



Wake of a fast boat
(surface wave analogy – dispersive)

See “Ship Wakes: Kelvin or Mach Angle?”
Rabaud & Moisy (PRL 2013)

Teleseismic observations of rupture speed

2. Surface wave Mach cone

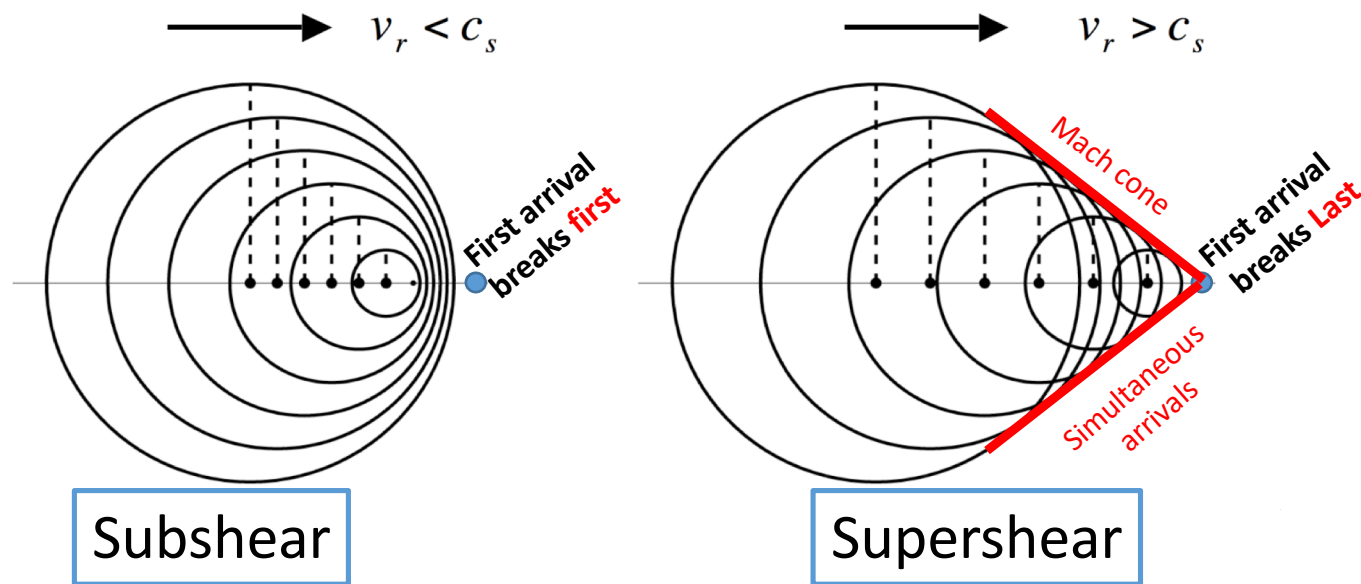
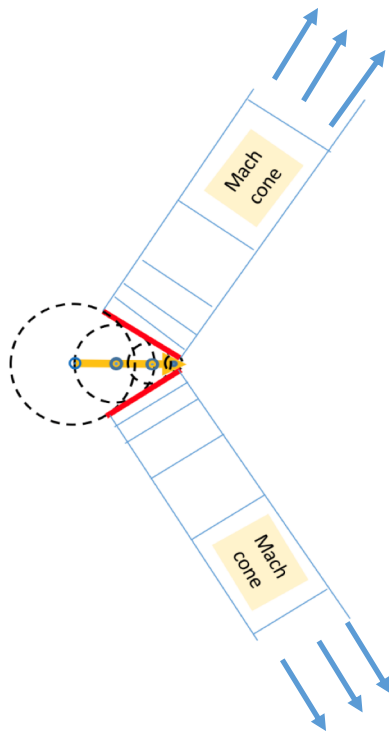


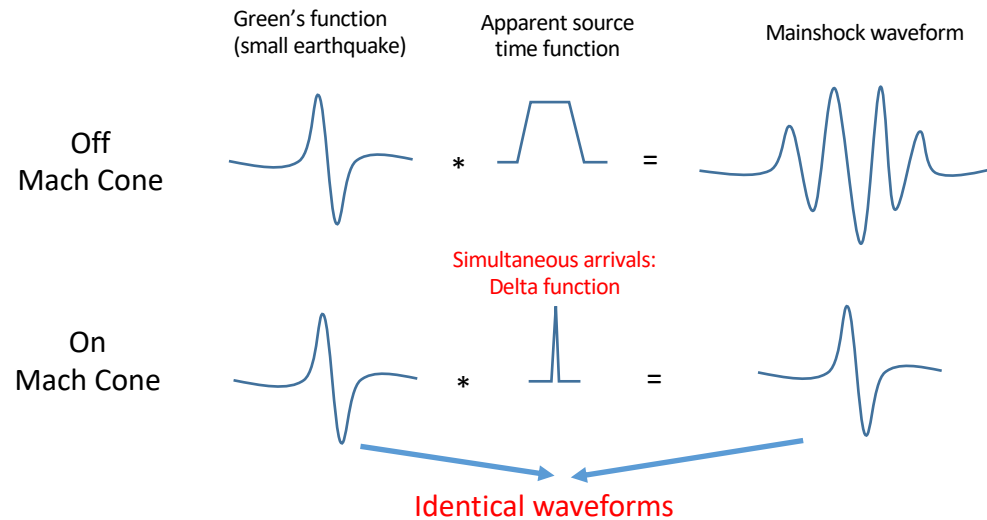
Figure modified from Eric Dunham's website

Teleseismic observations of rupture speed

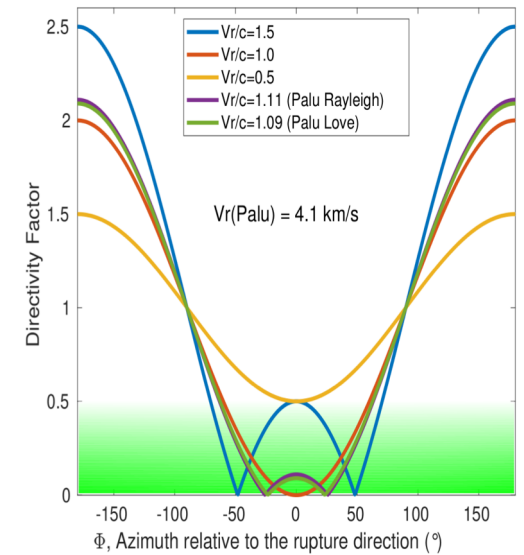
2. Surface wave Mach cone



WAVEFORM SIMILARITY OF SURFACE WAVE



apparent duration
real duration

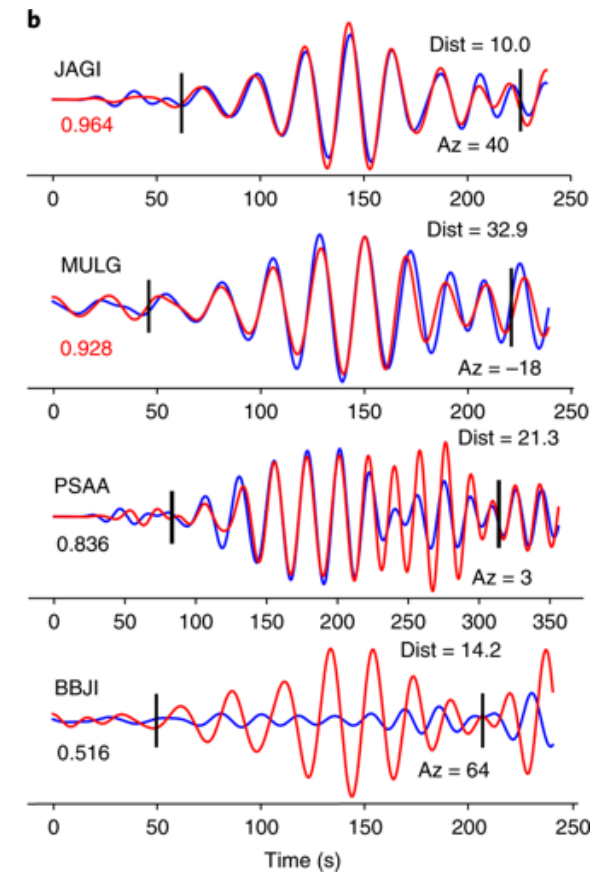
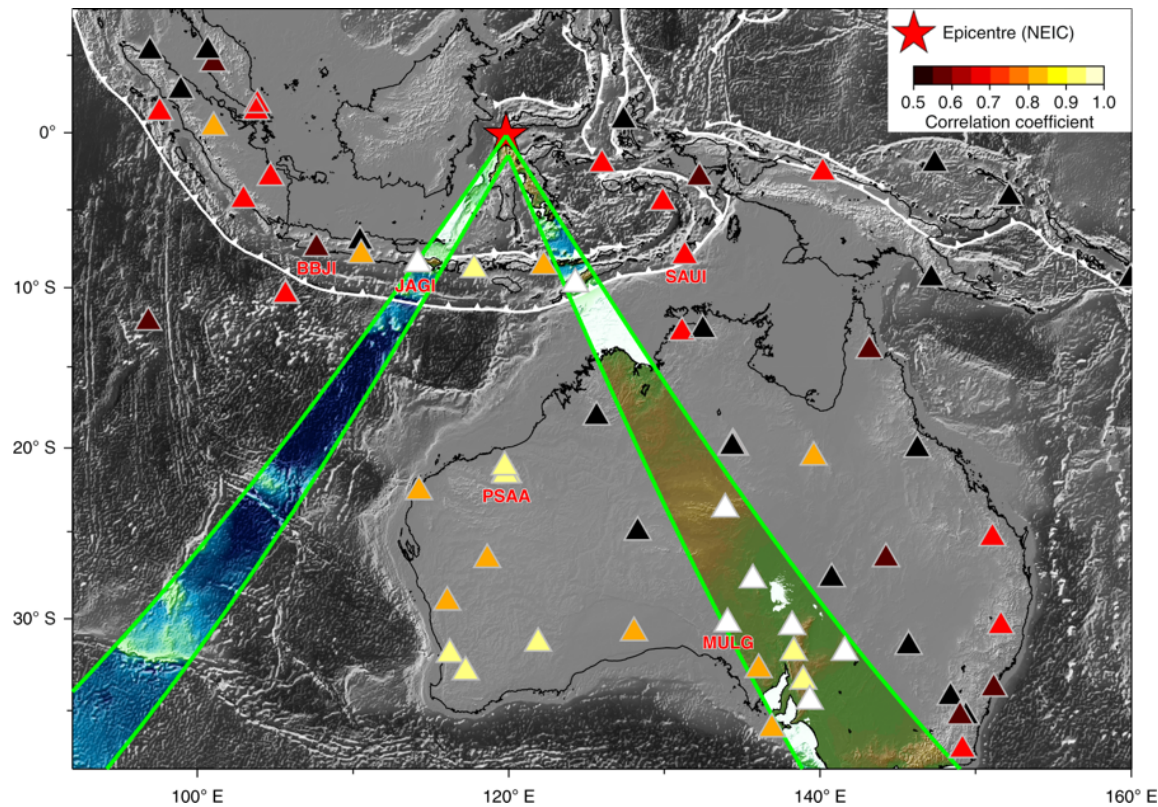


wave period
rupture duration

Following Vallée and Dunham, 2012

Teleseismic observations of rupture speed

2. Surface wave Mach cone



In the
Mach
cone

Off the
Mach
cone

Supershear earthquakes

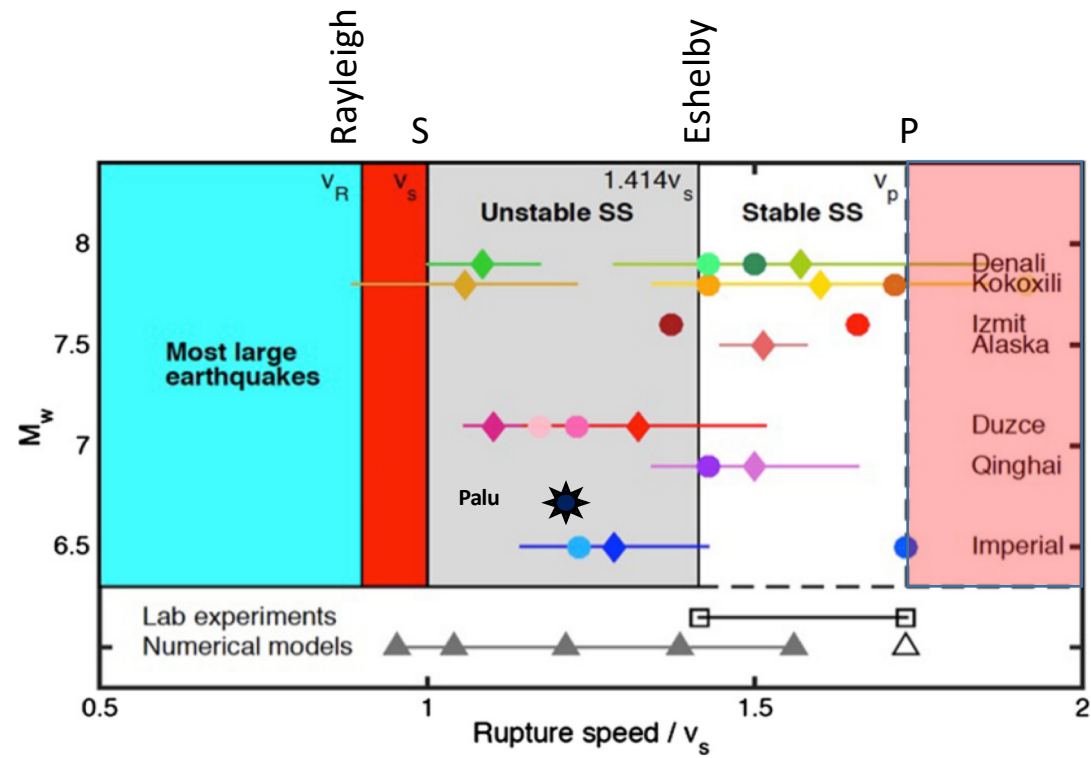


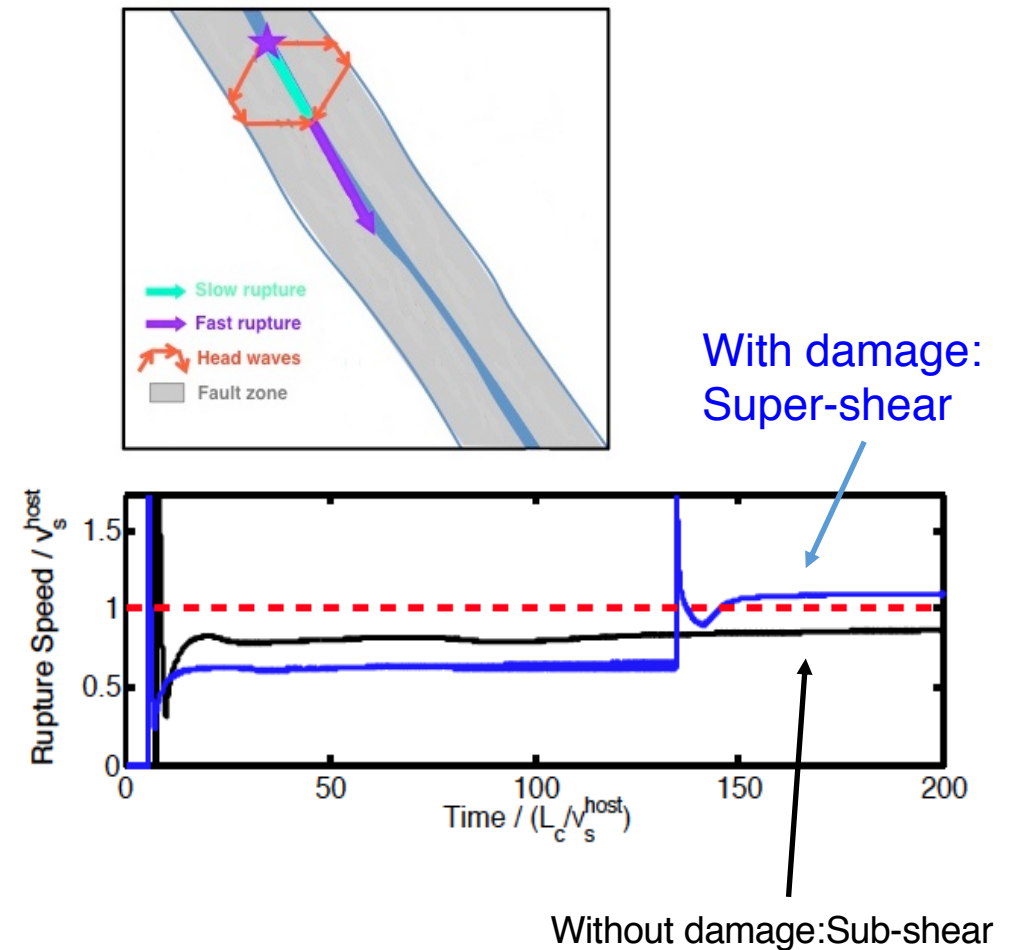
Figure modified from Huang et al., 2016

Fault zone waves facilitate supershear rupture

From dynamic rupture models:

The presence of a damaged fault zone facilitates supershear rupture

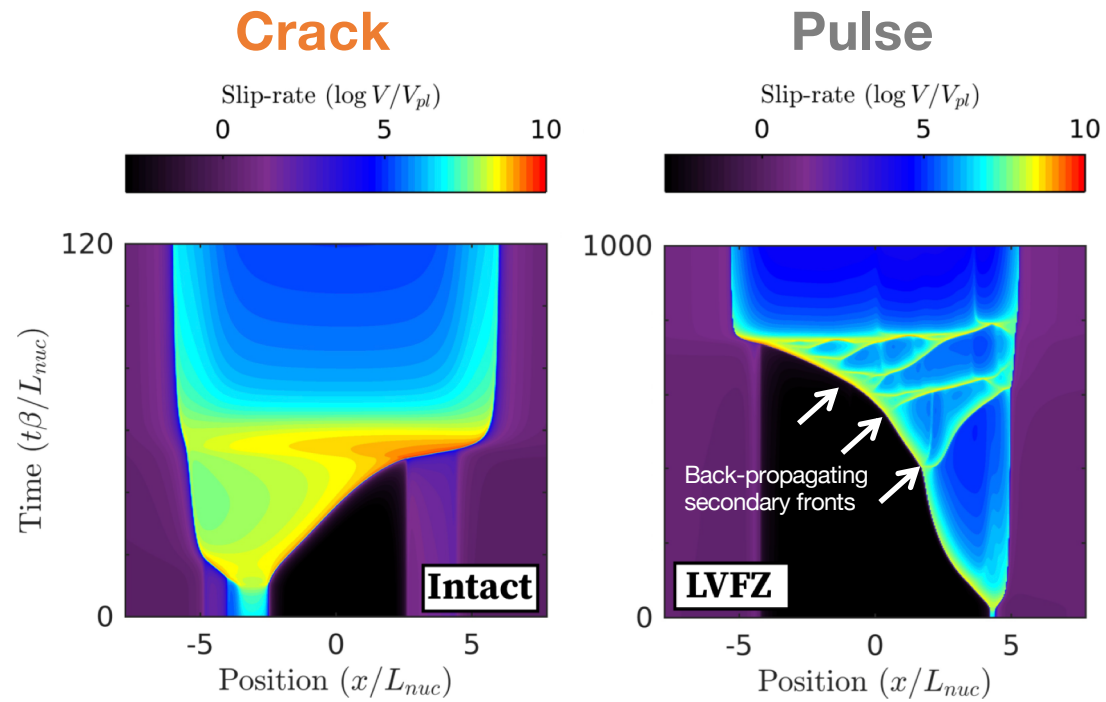
→ it can happen sooner and at lower stress than in intact rock



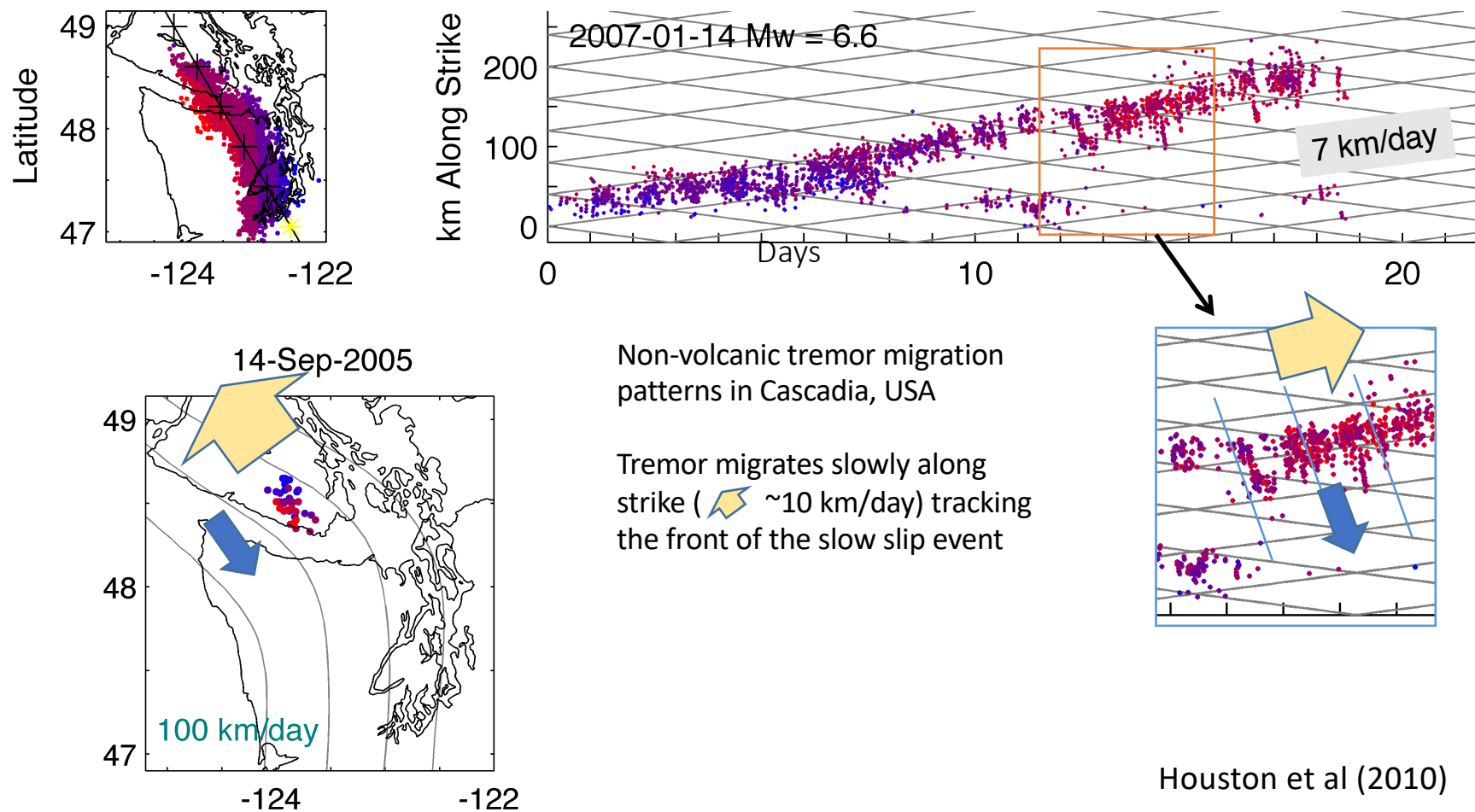
Huang and Ampuero (2011)

Huang, Ampuero and Helmberger (2014, 2015)

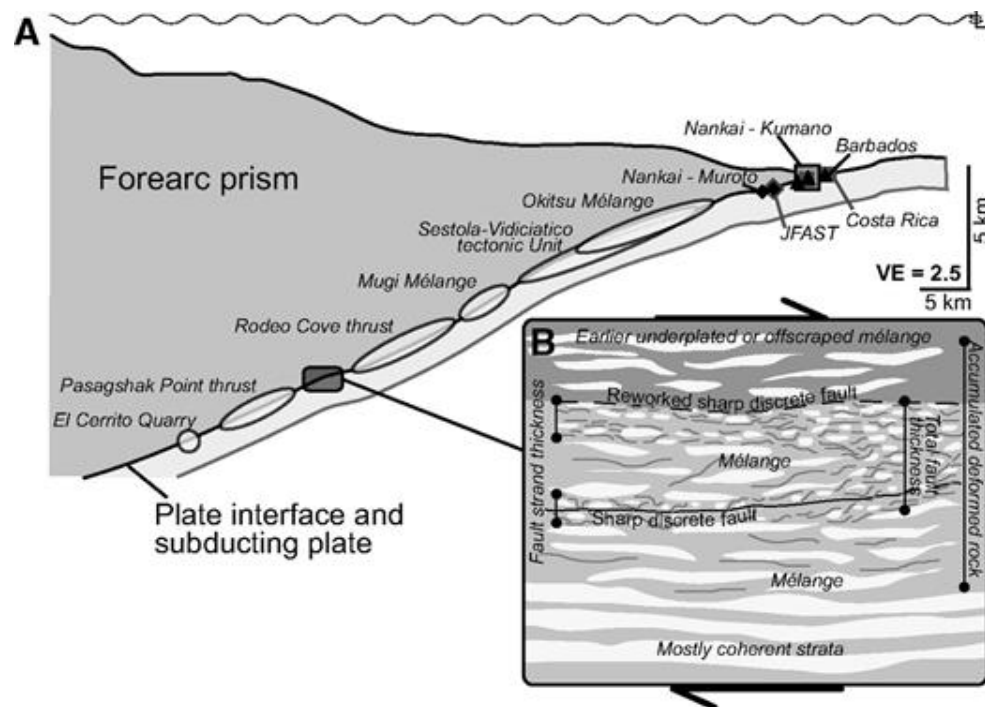
Back-propagating rupture pulses



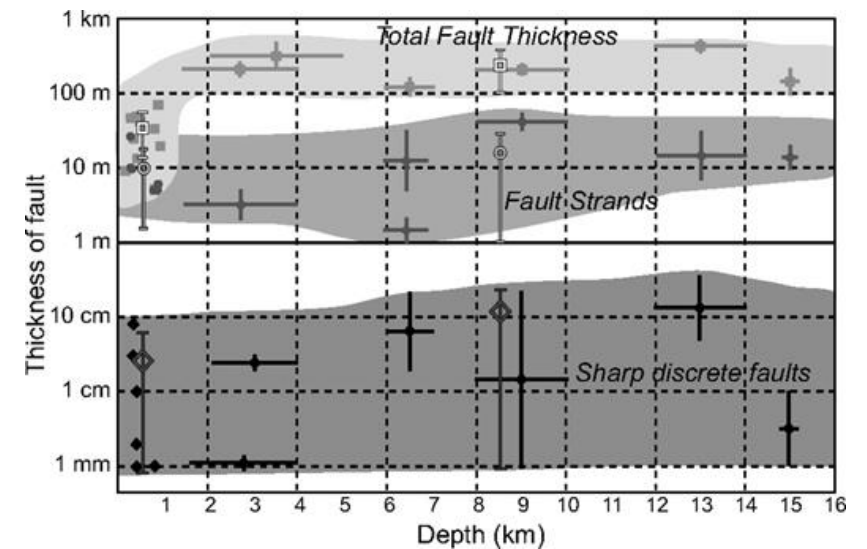
Rapid tremor reversals



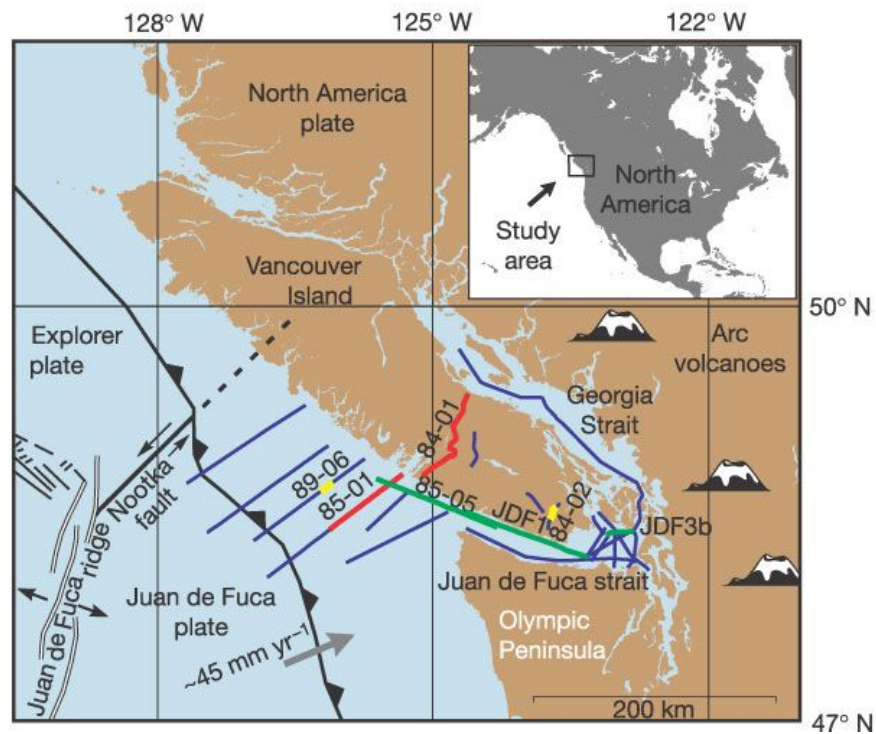
Low velocity zones in subduction megathrusts



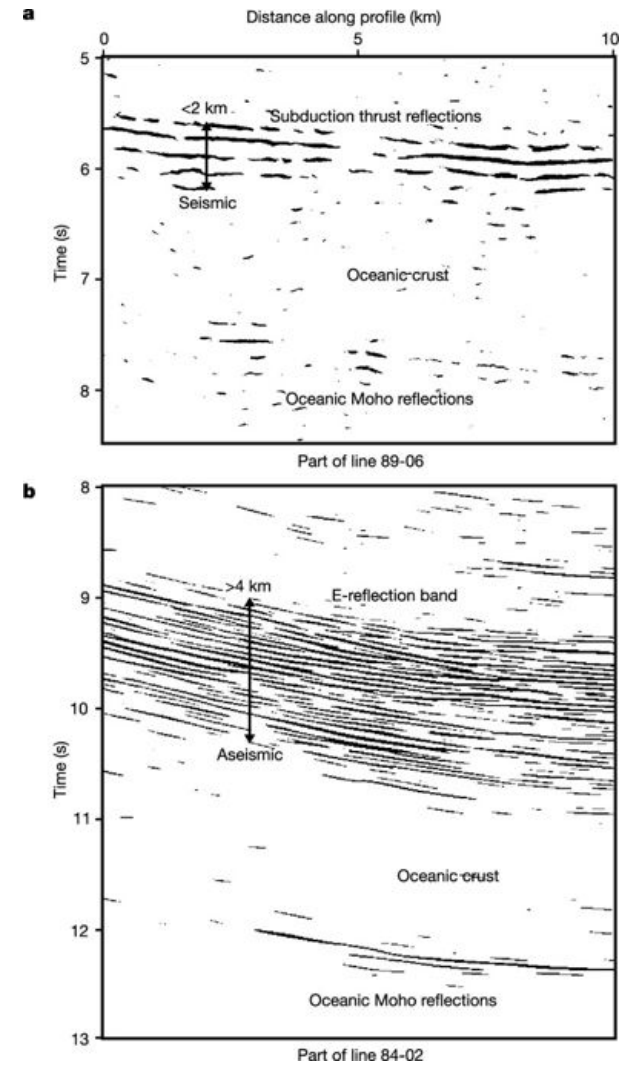
Rowe et al (2013)



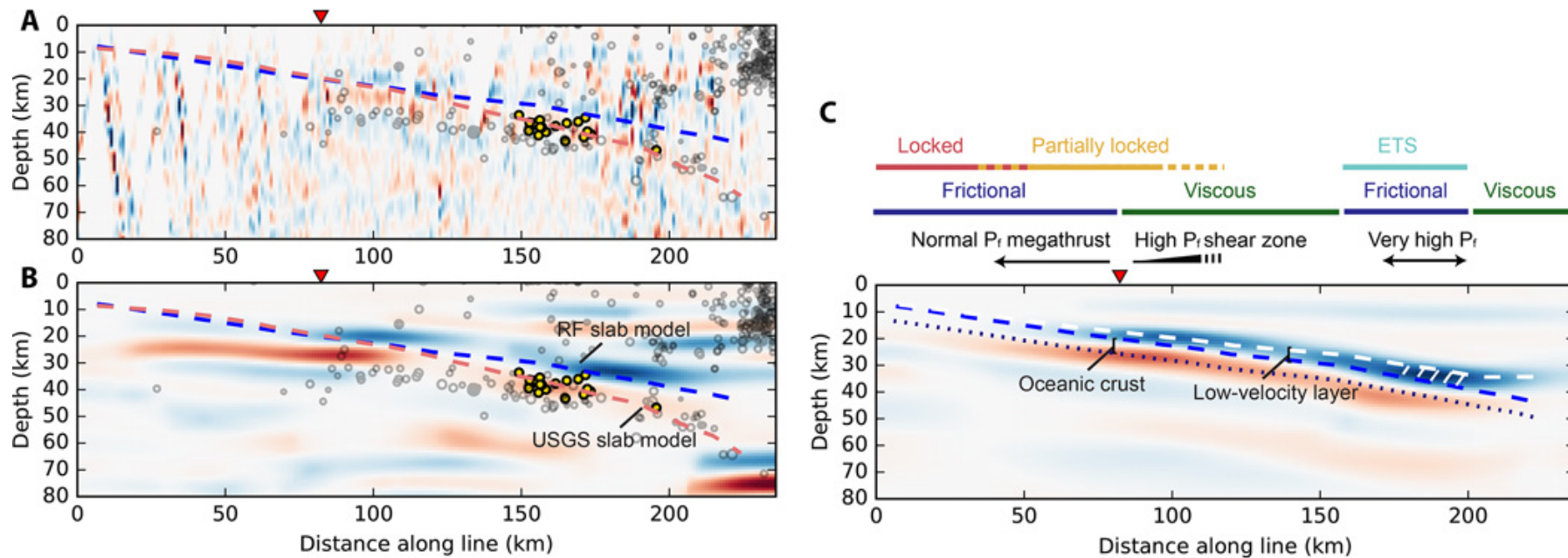
Low velocity zones in subduction megathrusts



Nemidovic et al (2003)

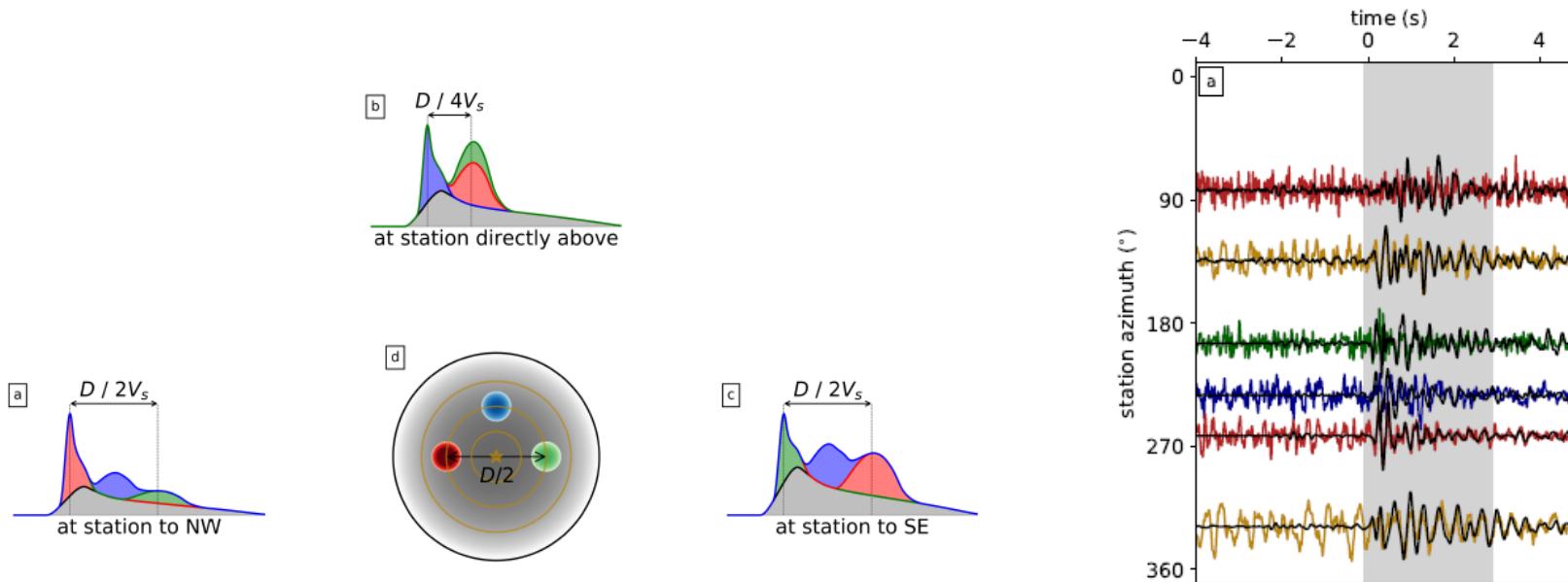


Low velocity zones in subduction megathrusts



Receiver function imaging of low velocity zones at tremor and SSE depths in Cascadia (Audet and Schaeffer 2017)

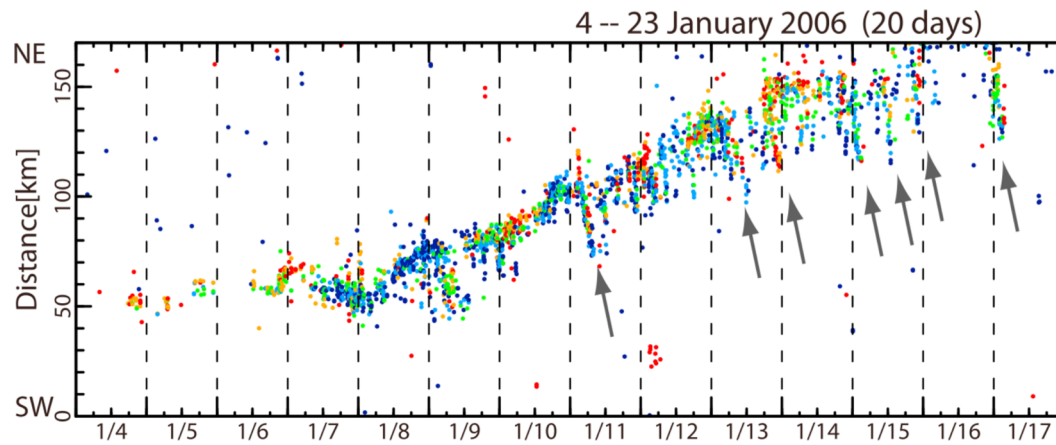
New constraints on LFE rupture speed (Hawthorne, Thomas and Ampuero 2018)



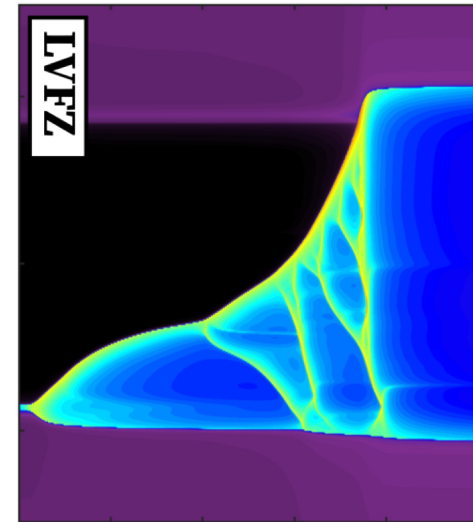
LFE rupture speed $< 40\%$ V_s of host rock

Back-propagating in slow slip model

Episodic tremor events in **Nankai, Japan**



Modified from Obara et al. (2012)



A potential candidate to explain rapid tremor reversals during SSE.