

**" Sixth Workshop on Non-Linear Dynamics and
Earthquake Prediction"**

15 - 27 October 2001

**The Dynamical Influence of Fluids
on Seismicity and Faulting**

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Lecture I: The dynamical influence of fluids on seismicity and faulting

- Rock and fluid mechanics
- Dislocations and brittle faulting
- Stress transfer models
- The hydraulics of fault zones
- Coupling fluid flow to large scale tectonics
- Earthquakes as a coupled shear stress, high pore pressure dynamical system
- The properties of large model earthquakes

Lecture II: Earthquake scaling and the strength of seismogenic faults

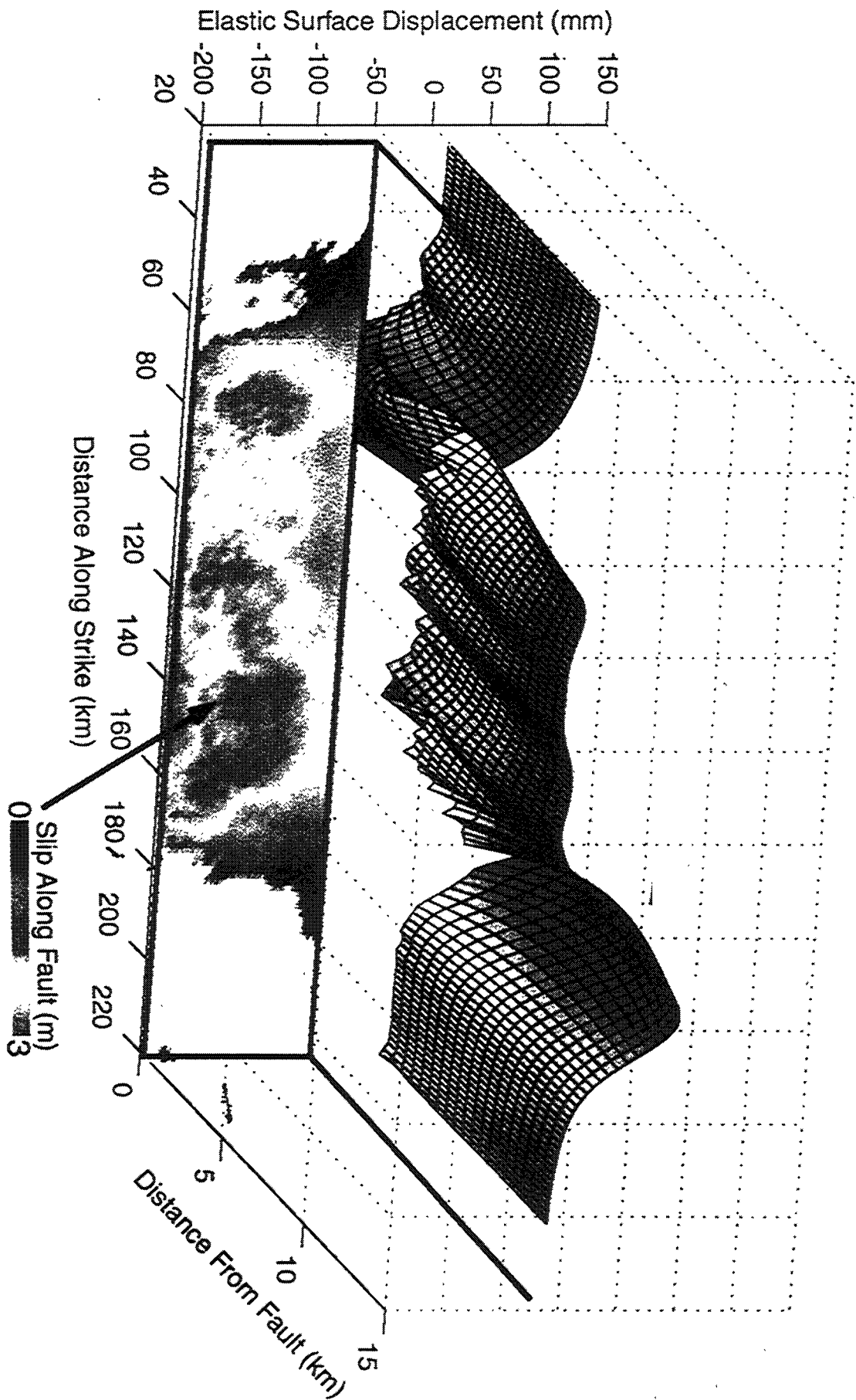
- Earthquake scaling: Small earthquakes
- Earthquake scaling: Large earthquakes
- Pore pressure as a fundamental scaling parameter
- Application of result to global earthquake catalogs
- Can pore pressures be inferred from earthquake rupture properties

Lecture III: Neotectonics and earthquake generation along complex fault systems

- The North Anatolian Fault Zone
- The San Andreas Fault System
- Building a forward model of the earthquake process
- Fault Interaction

The Model

- Fault-valve behavior and fault zone architecture
- Permeability as a toggle-switch
- Coupling fault zone pore pressures to large-scale tectonic loading
- Stress space disorder
- Physical space disorder
- Slip weakening friction
- Surface displacements in response to slip
- Long-term fault behavior
- 1. Stress Drop
- 2. Average Slip vs. Rupture Length
- 3. Comparison with earthquake catalogs
- Generalizing the model to any geometry



Strength of the Lithosphere

Frictional Heating

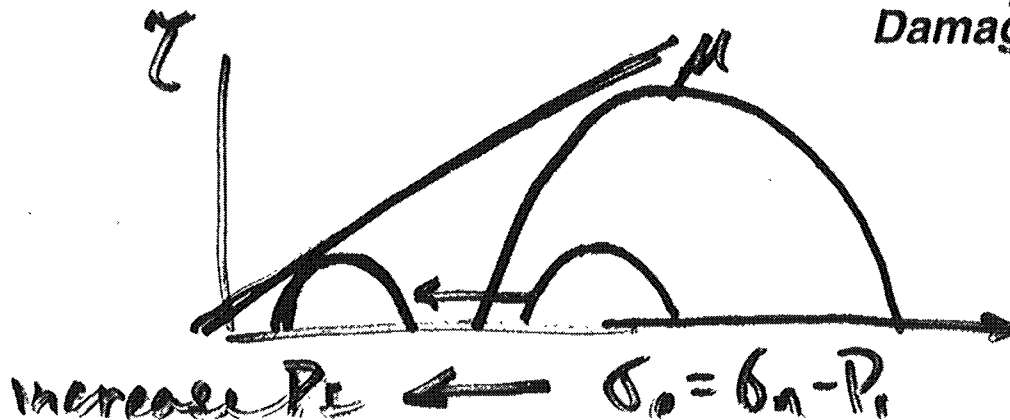
*Dehydration
Melt
Afterslip*

Fault Strength
(Pore Pressure)

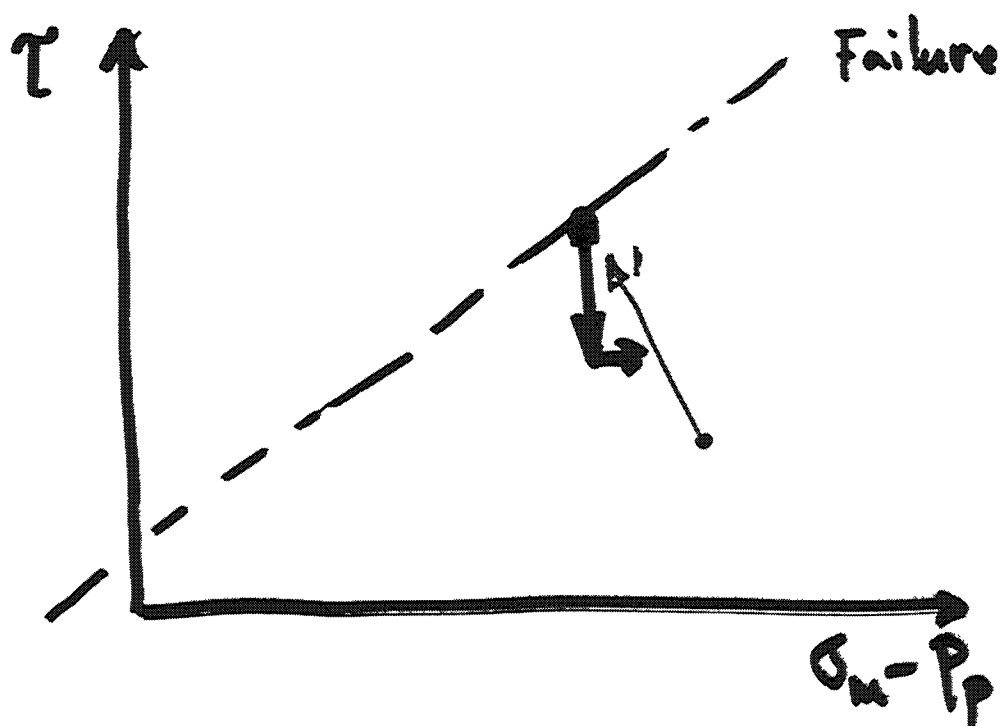
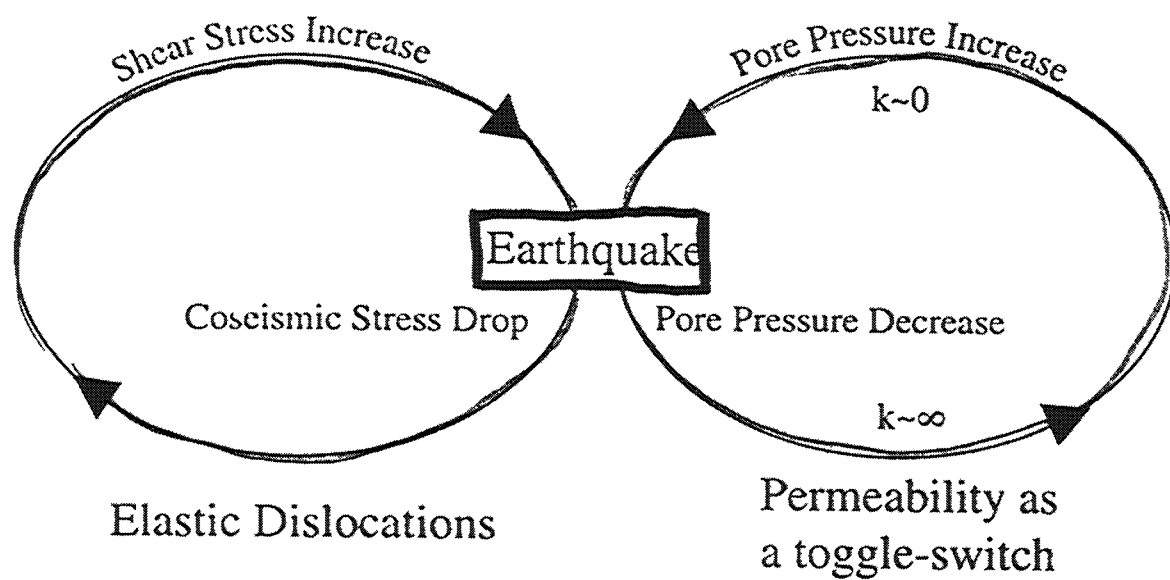
Plate Kinematics

Earthquake Magnitude

*Rupture Length
Damage (permeability)*



Fault Valve Model (Sibson, 1992)



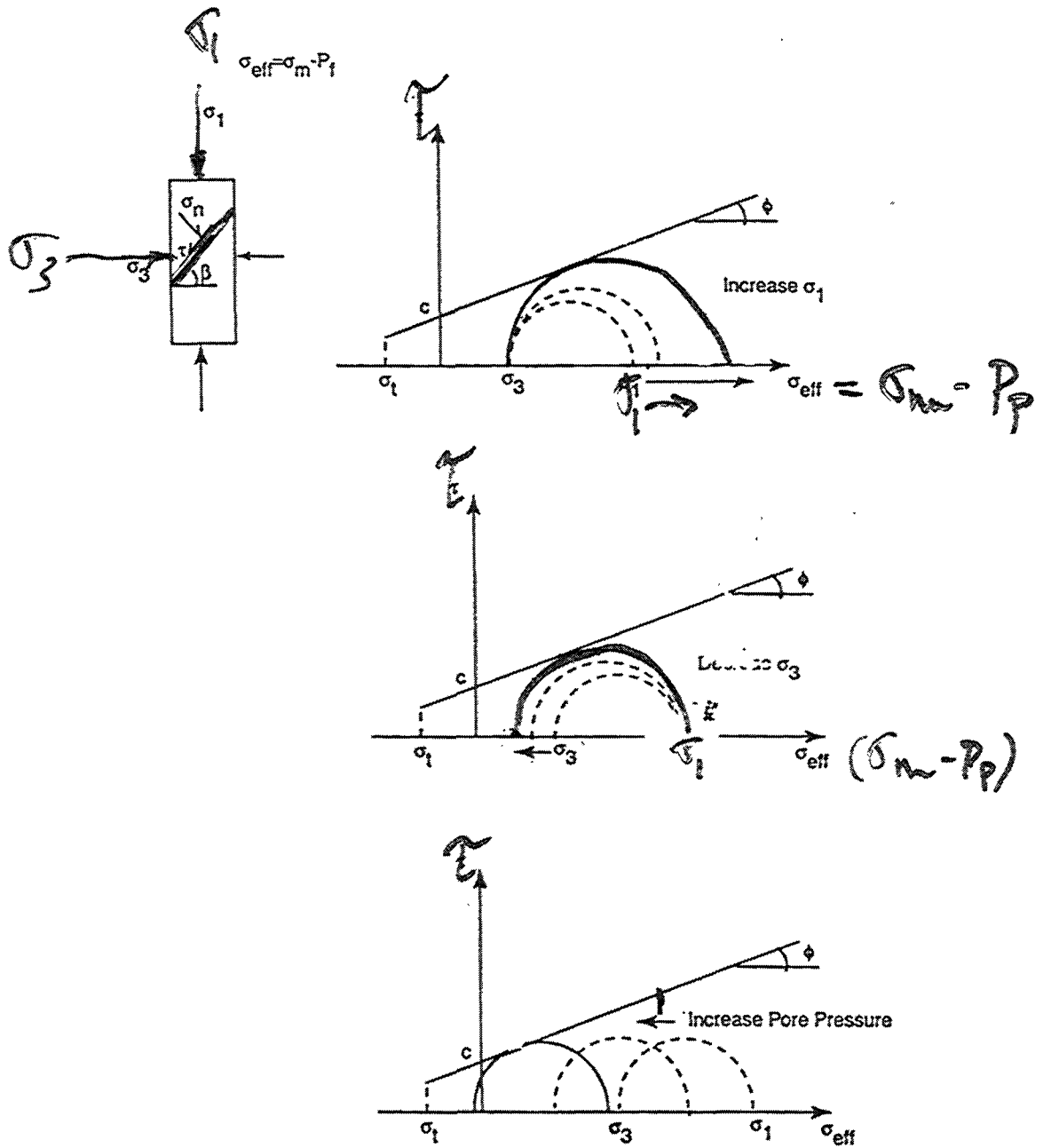
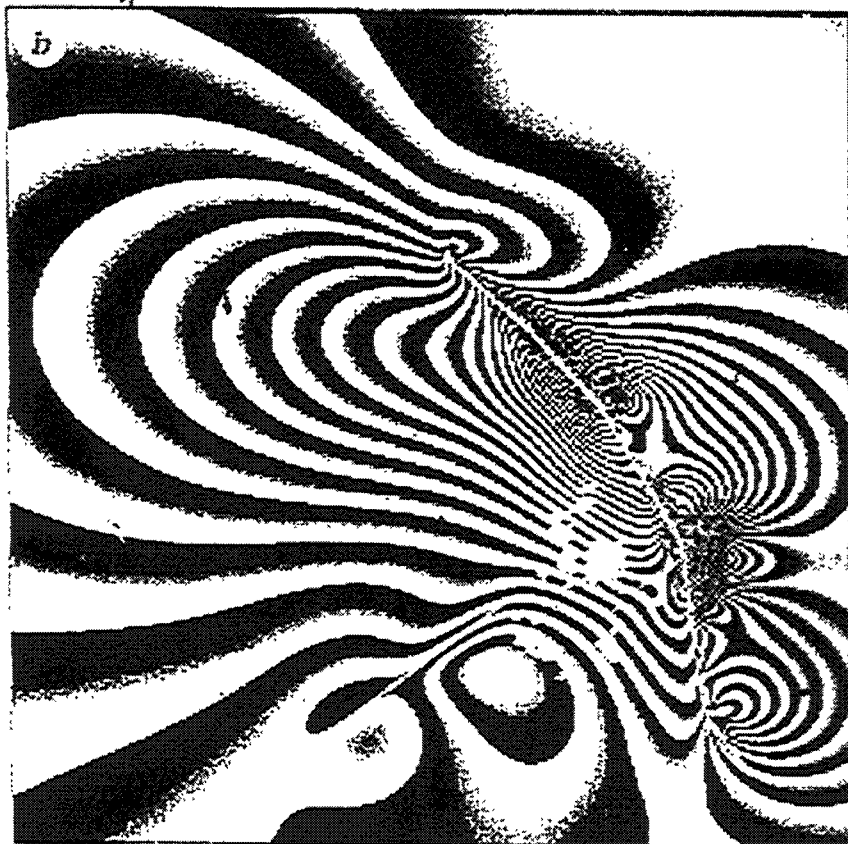
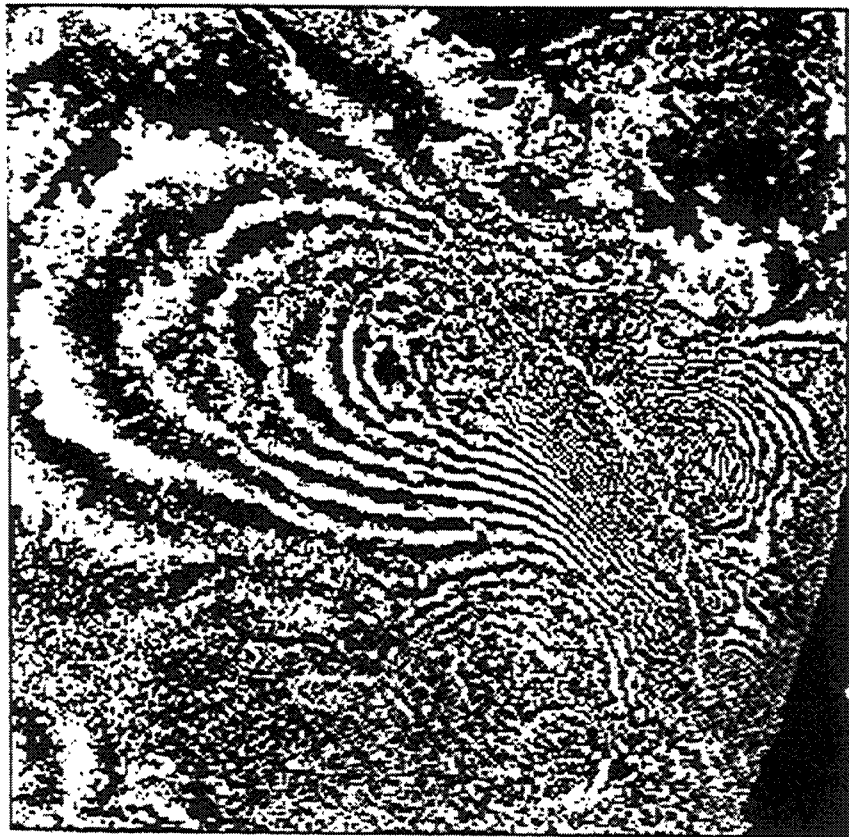


Figure 3.11: The influence of fluids on rock strength

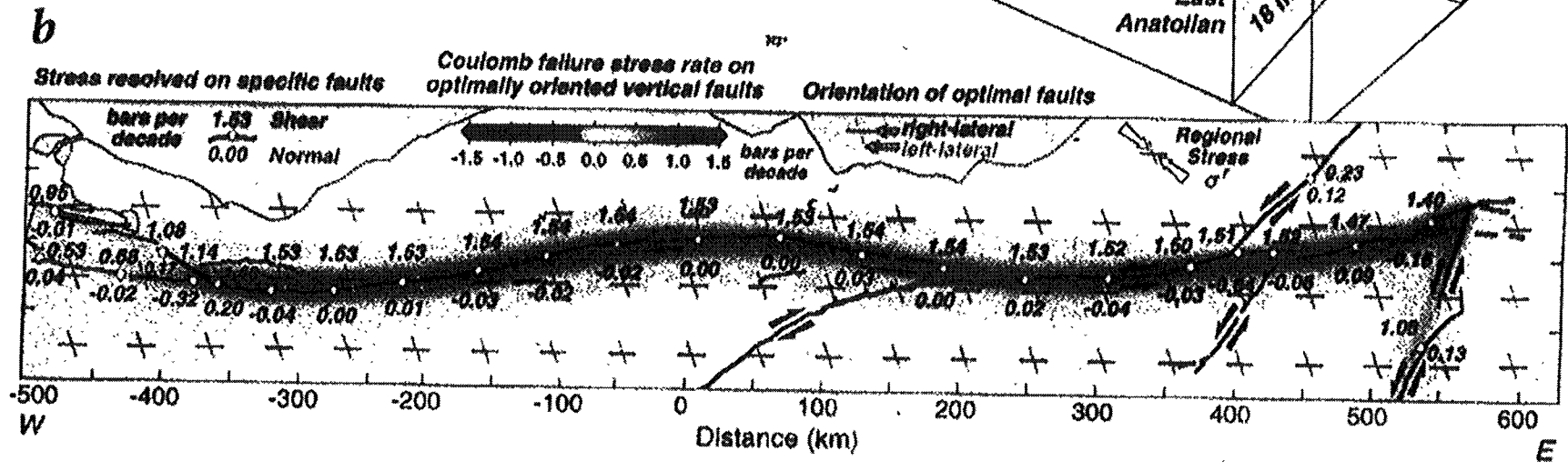
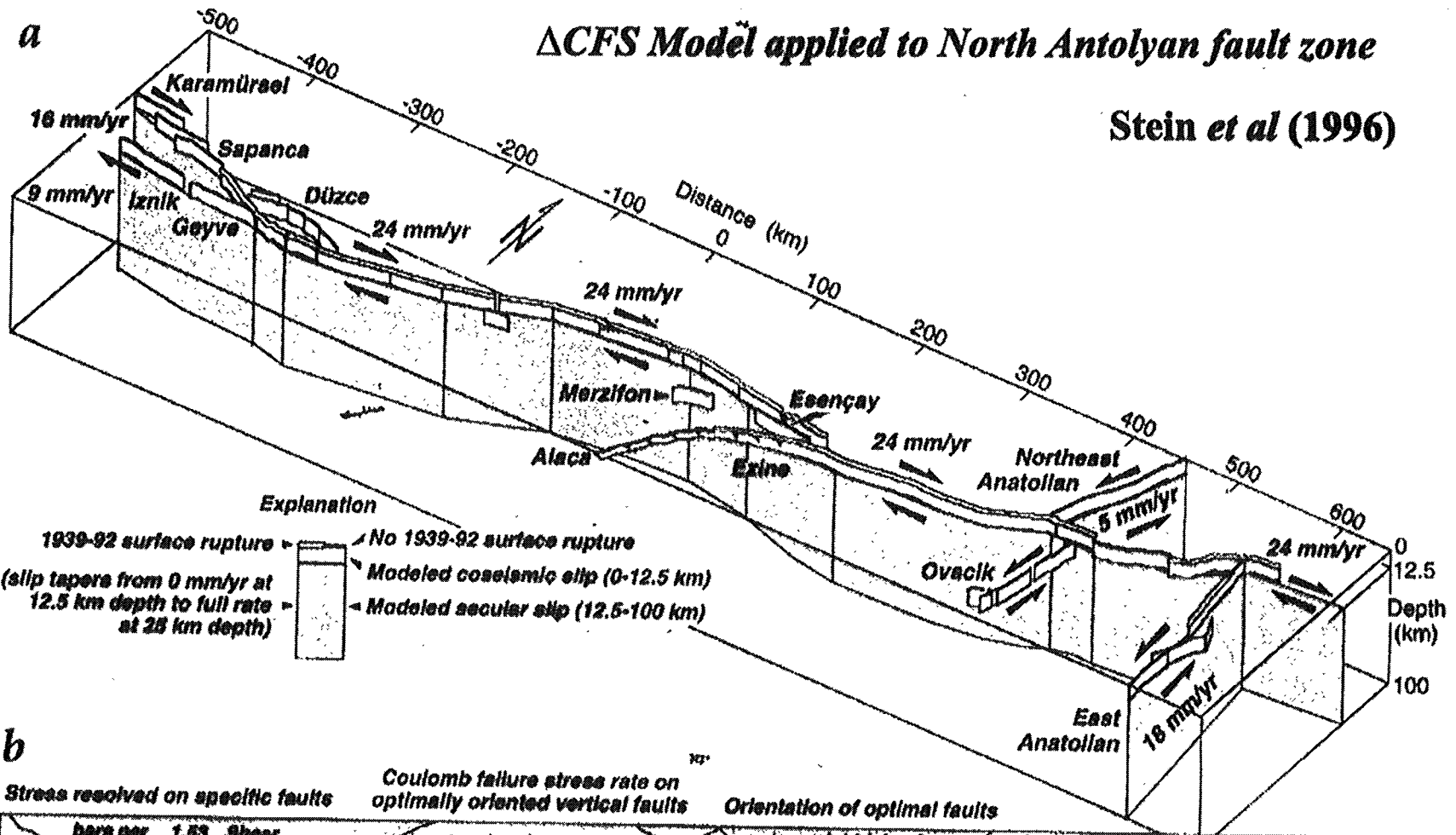
Using Elasticity Theory to Monitor Analyze Earth Deformations



Massonne et al
(1993)

Δ CFS Model applied to North Antolian fault zone

Stein *et al* (1996)



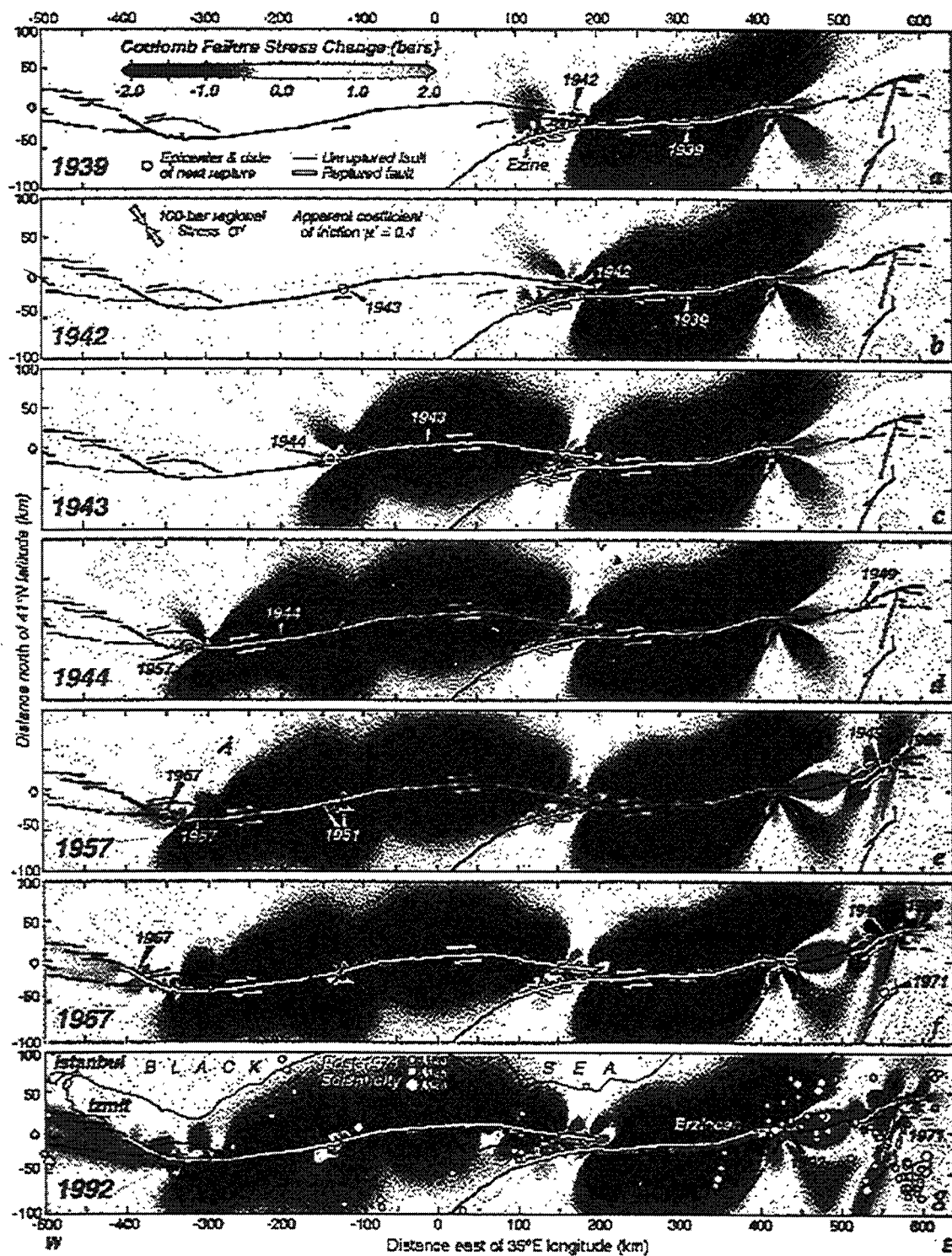
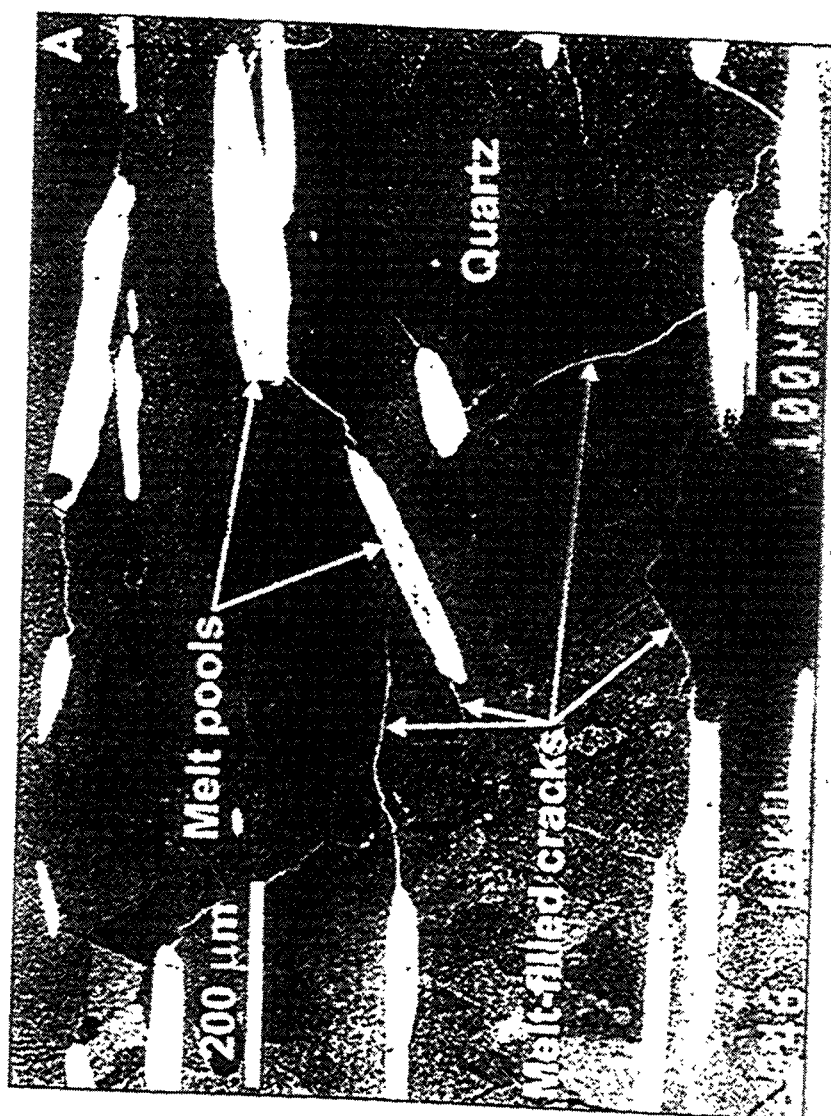
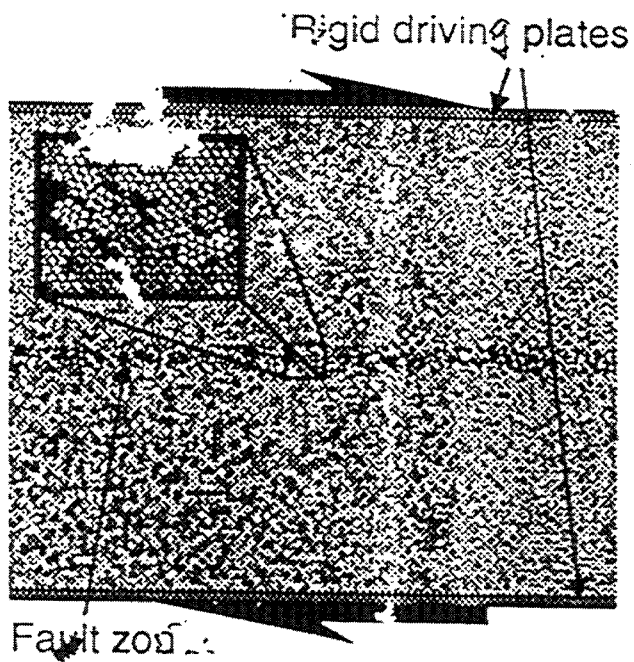


Figure 4 17 Oct 96 Stein et al.

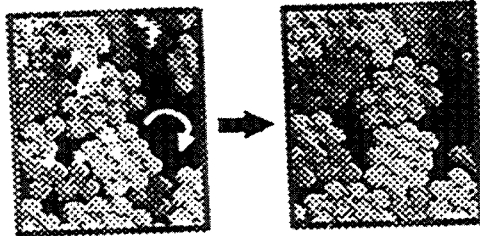


Cannary et al. (97)

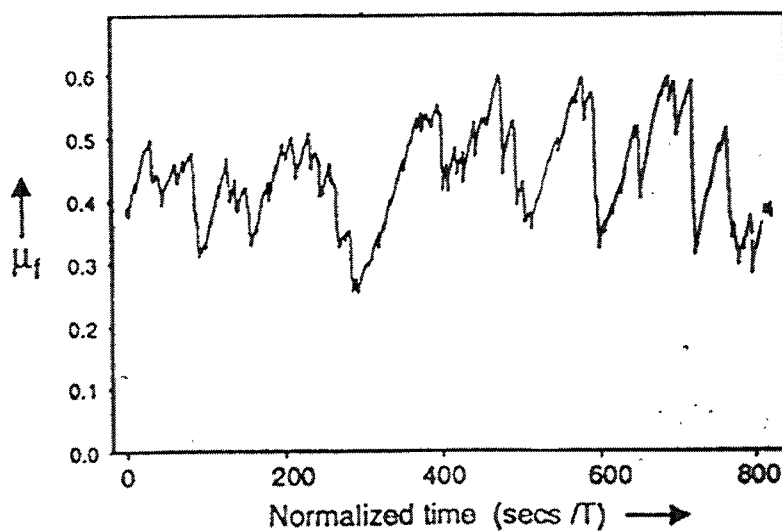
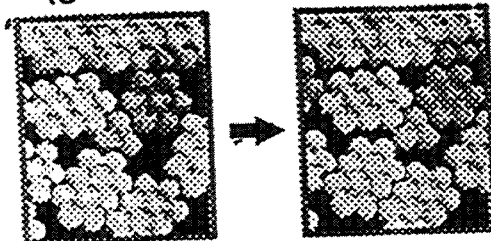


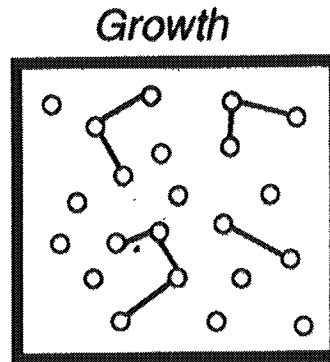
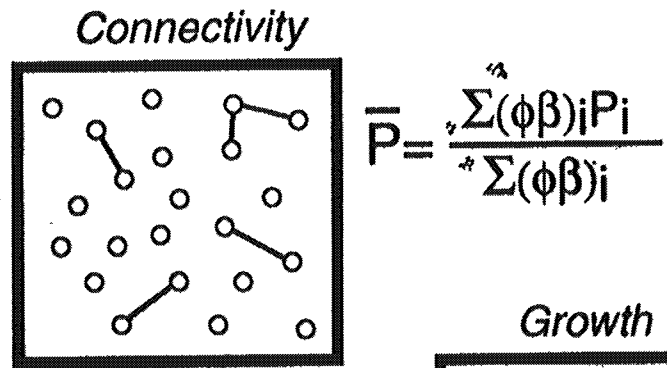
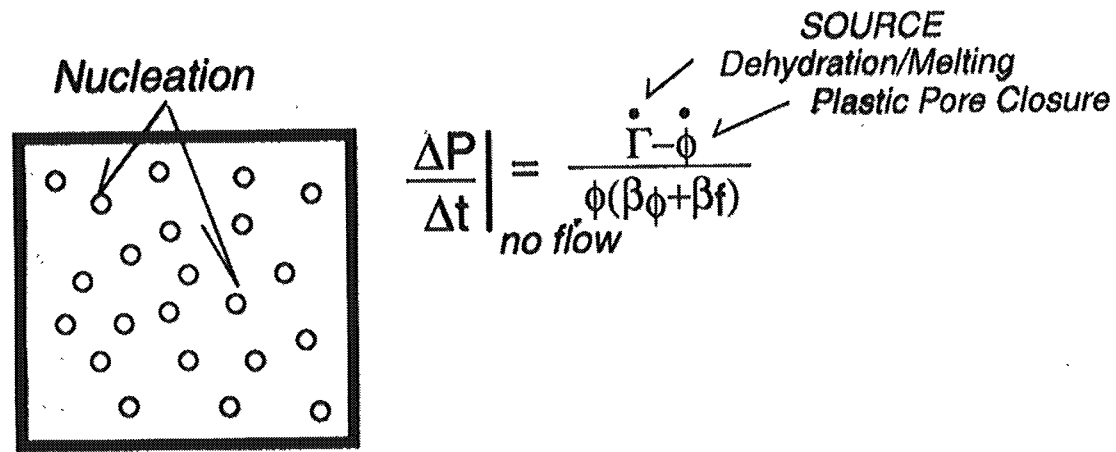
Place & Move

Weak fault
(grains roll during slip)

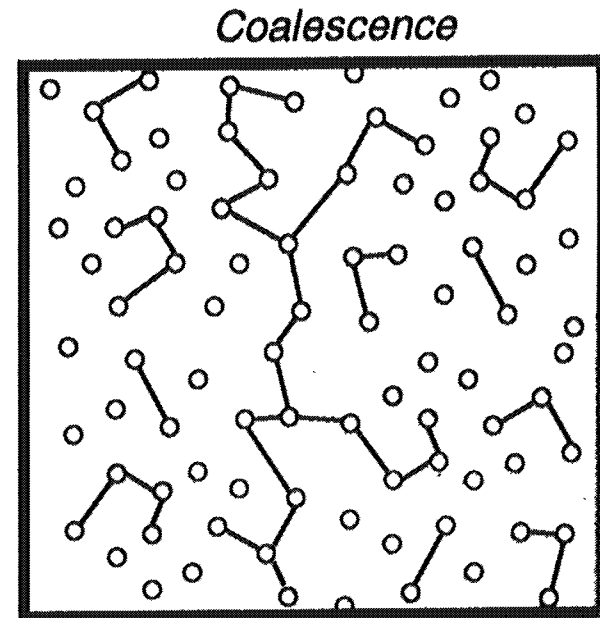
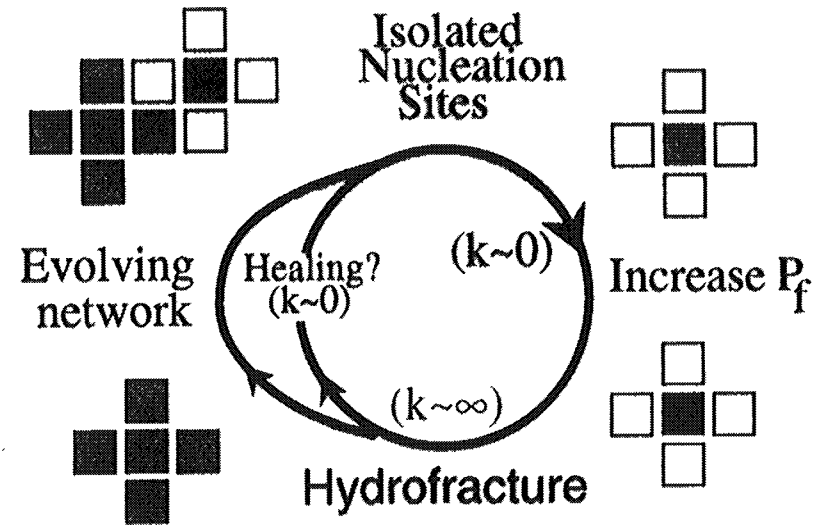


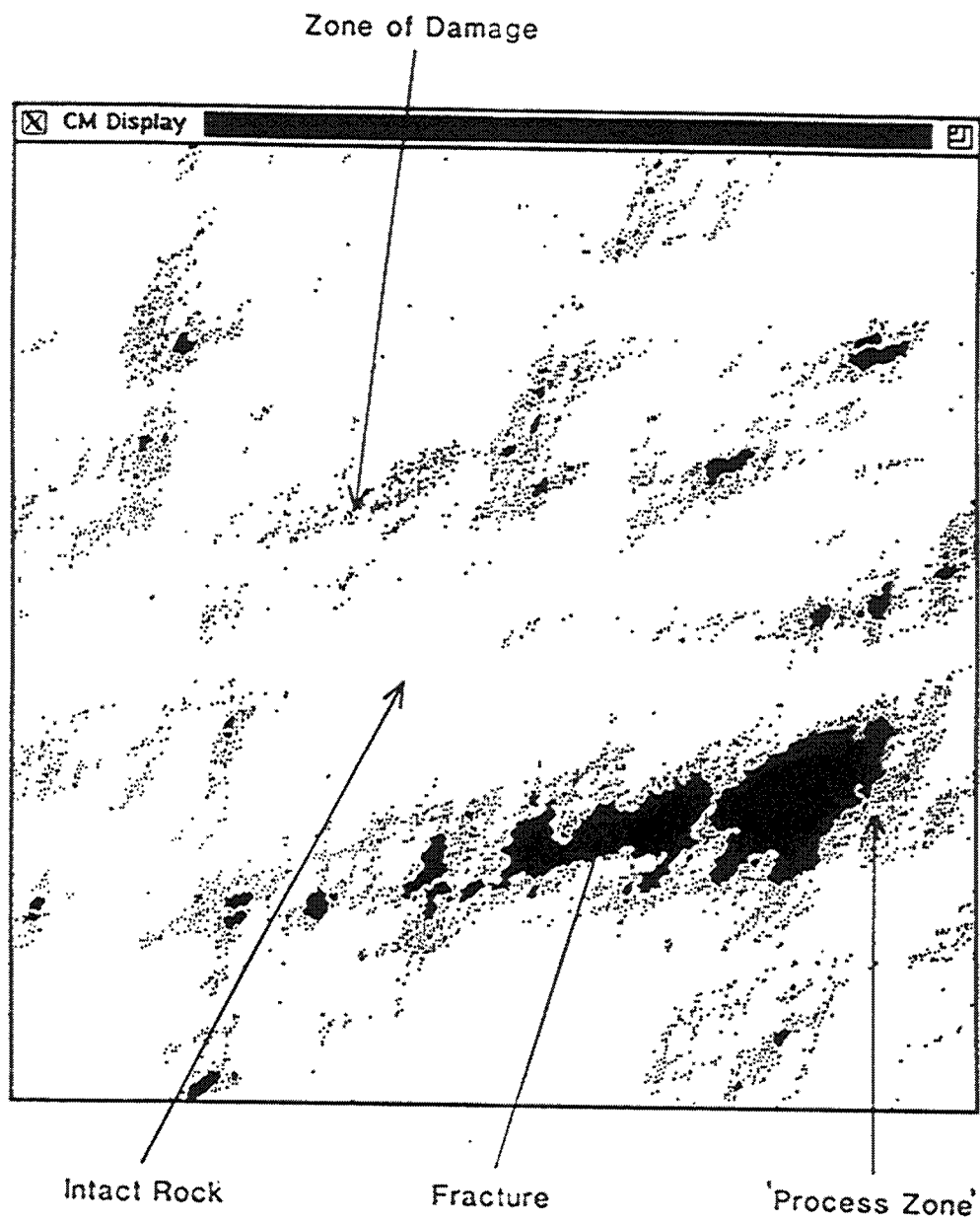
Strong fault
(grains slide during slip)





Conceptual Model





Henderson et al (1997)

Diffusion Equation with Source Term

$$\frac{\partial P_f}{\partial t} = \frac{1}{\phi(\beta_\phi + \beta_f)} \left[\frac{k}{v} \nabla^2 P_f - (\dot{\phi}_{plastic} - \dot{\Gamma}) \right] \quad (1)$$

where

- ϕ is porosity
- β_ϕ and β_f are the pore and fluid compressibility
- v is the viscosity
- ρ is the fluid density
- k is the intrinsic permeability of the matrix.
- $\dot{\phi}_{plastic} - \dot{\Gamma}$ is a source term

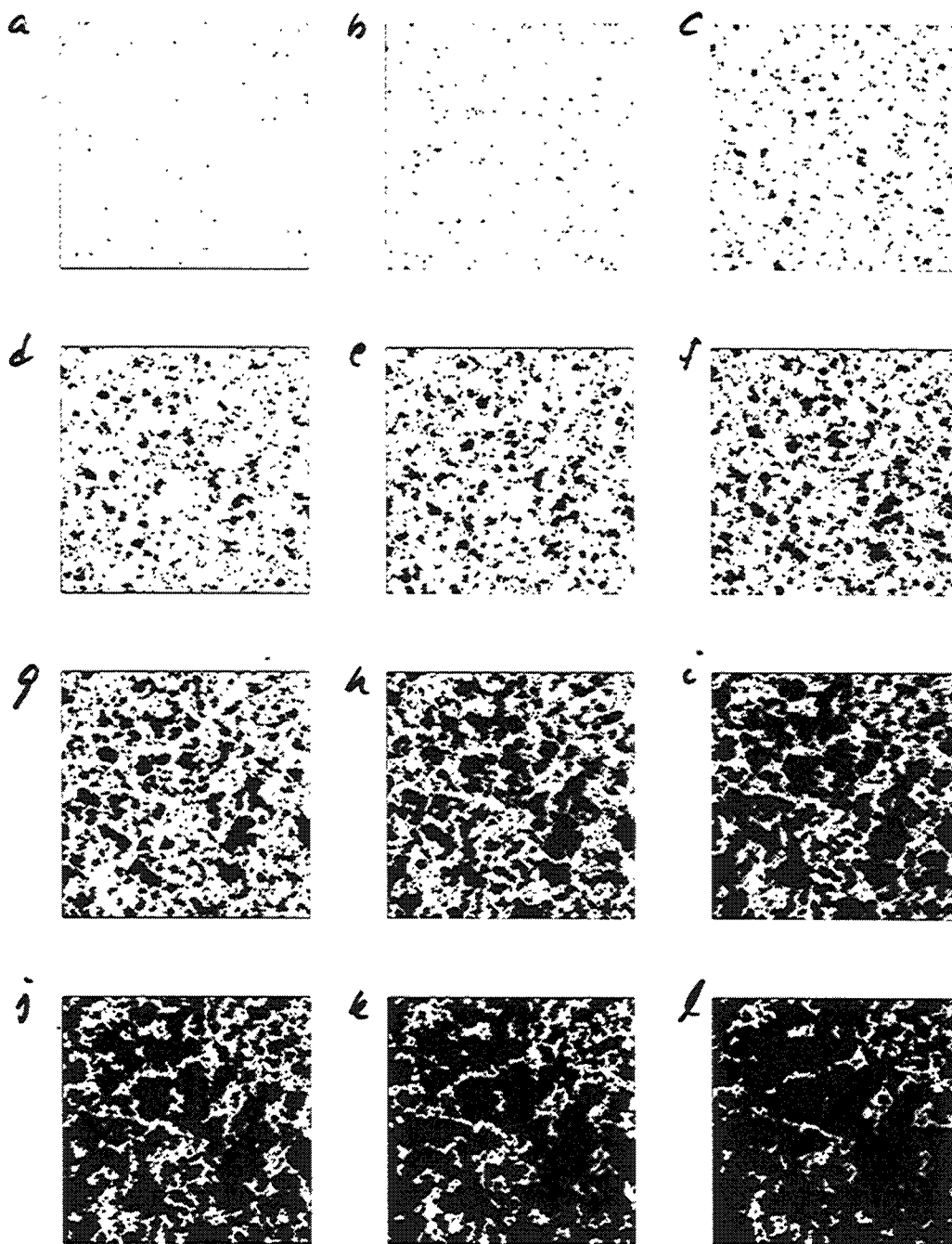
For an impermeable matrix ($k \sim 0$)

$$\frac{\partial P_f}{\partial t} \big|_{noflow} = \frac{(\dot{\Gamma} - \dot{\phi})_i}{\phi_i \beta_i} \quad (2)$$

At failure (hydrofracture), fluid pressures equilibrate with nearest neighbor cells. The equilibrium pressure (by conserving fluid mass and ignoring gravity, is:

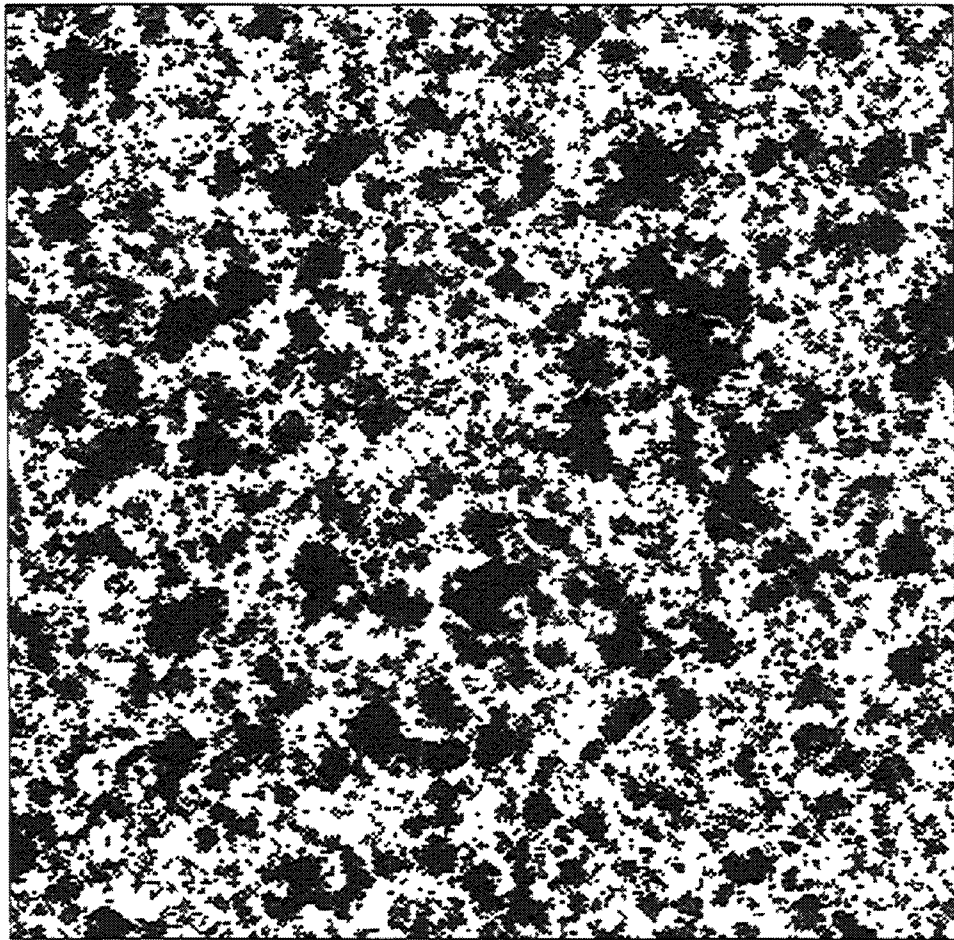
$$\bar{P} = \frac{\sum_{i=1}^m (\phi\beta)_i P_i}{\sum_{i=1}^m (\phi\beta)_i} \quad (3)$$

where $\bar{P}$ is the average pressures of affected cells.

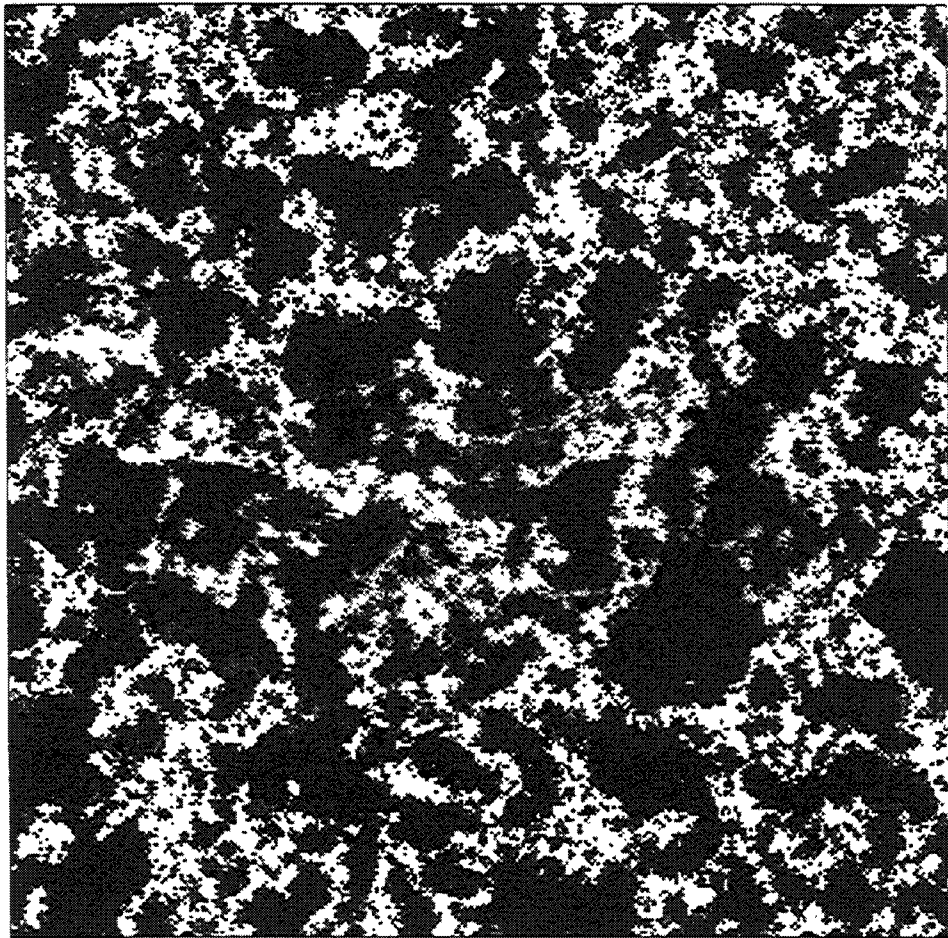


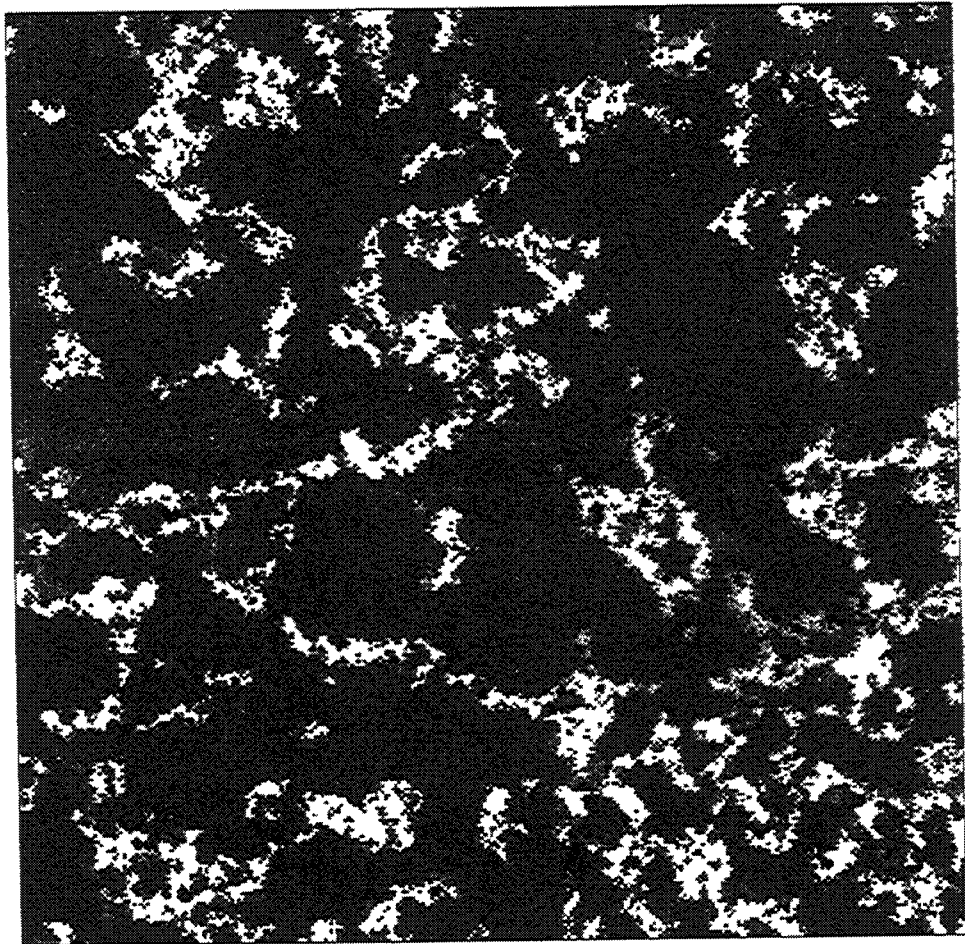
Black = Incipient Fashene

Miller & Naur (2000)

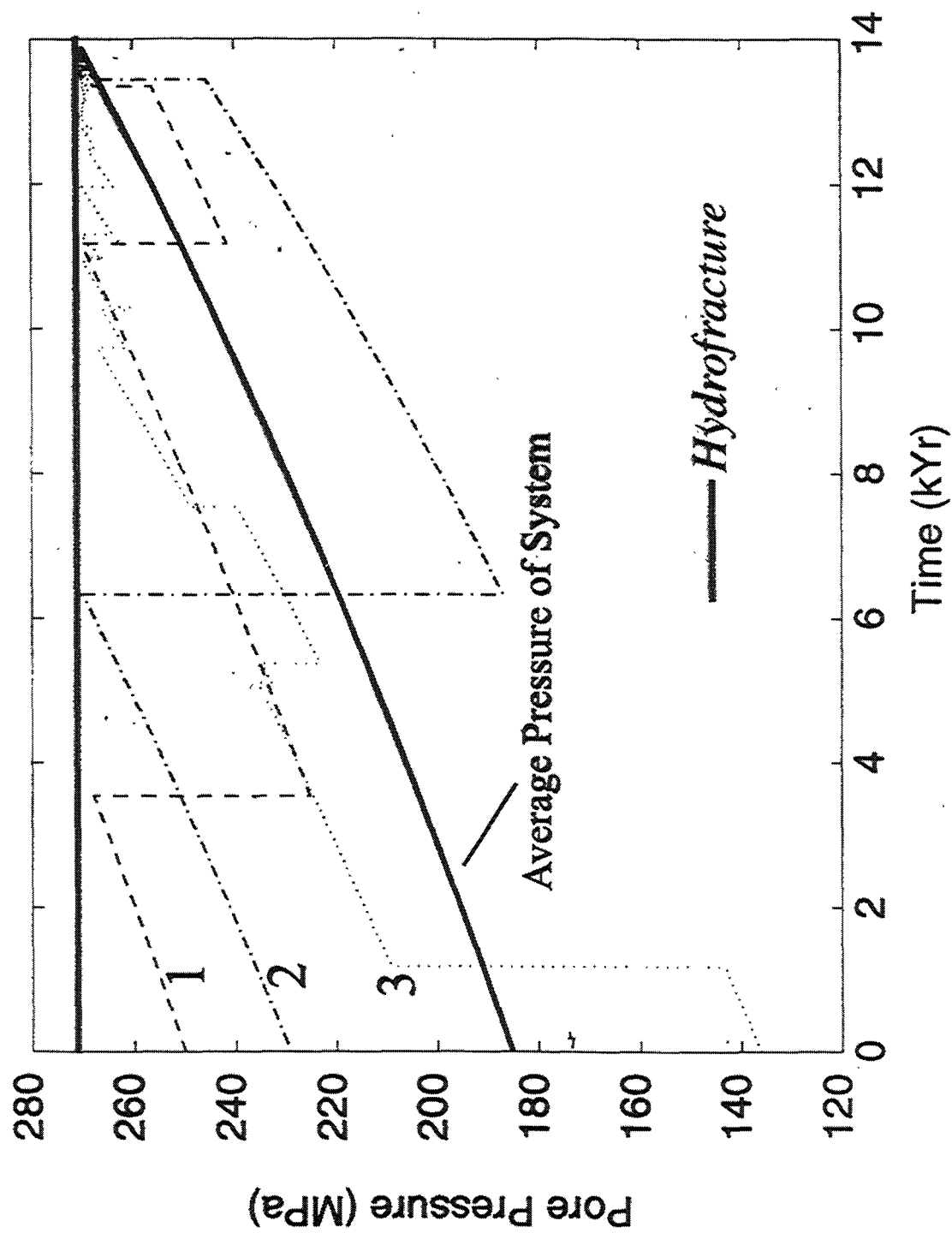


Cells at failure Condition

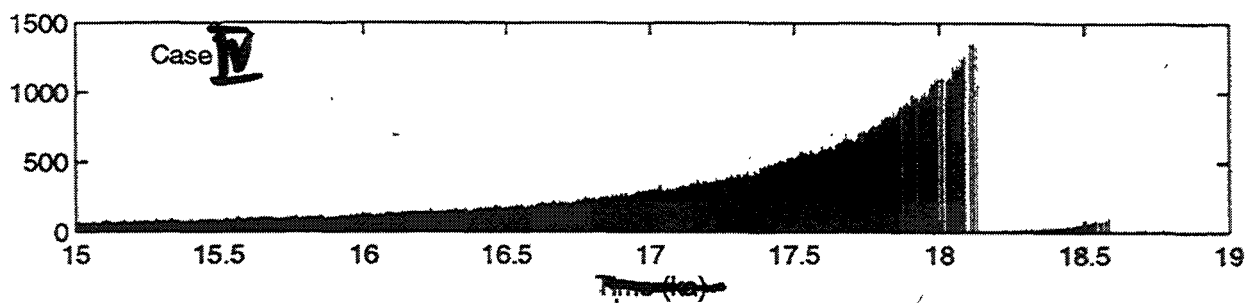
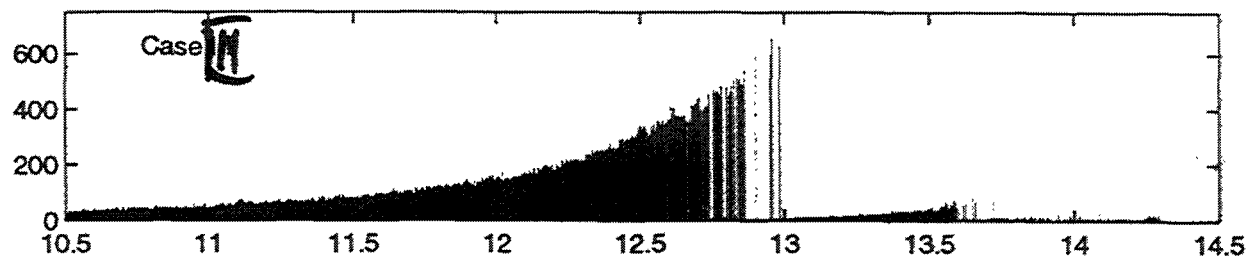
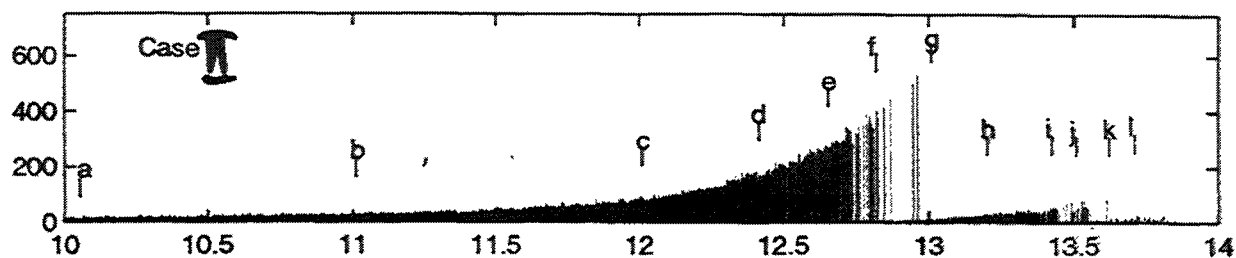
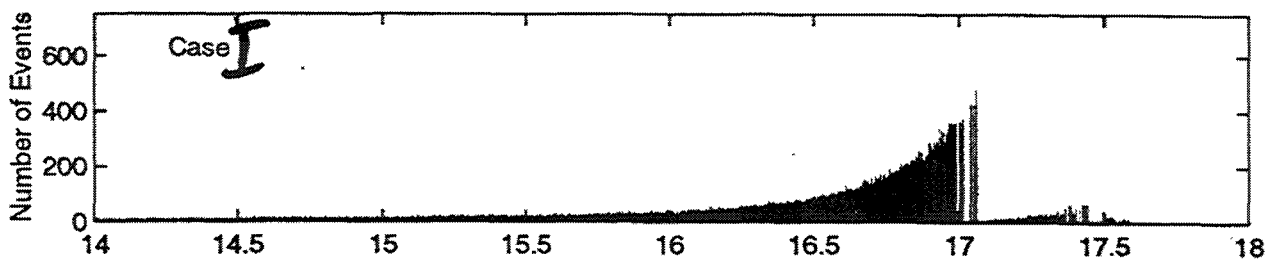




Pore pressure paths of arbitrary cells and system pressure



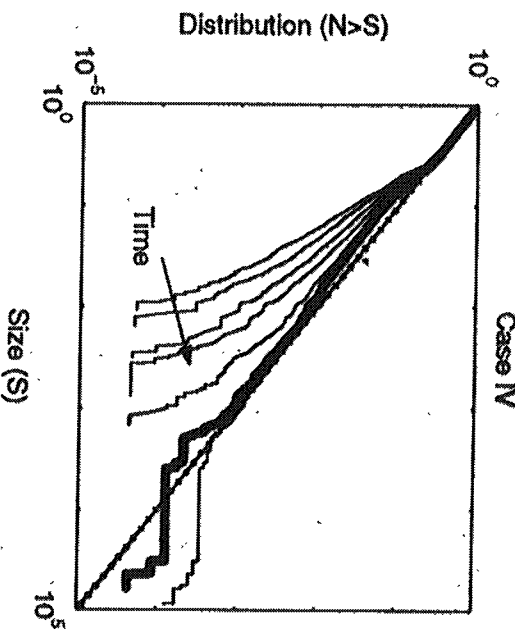
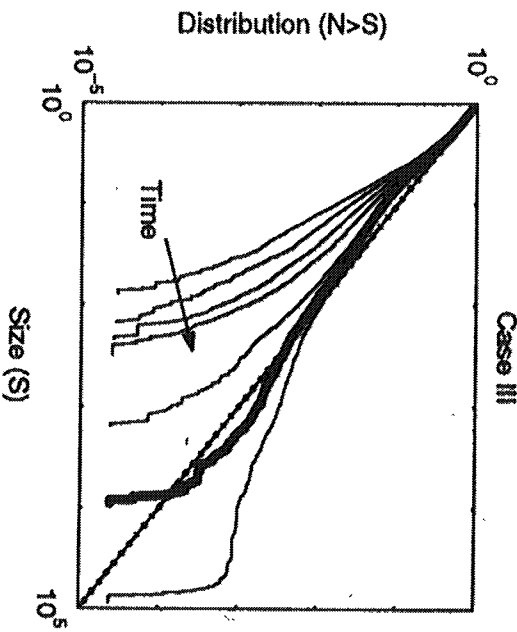
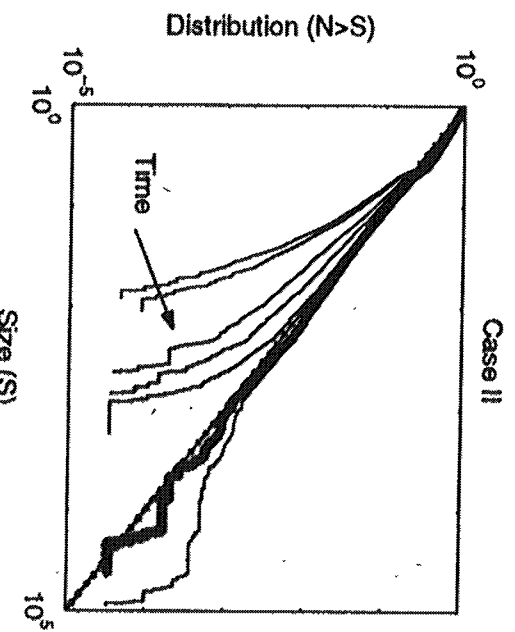
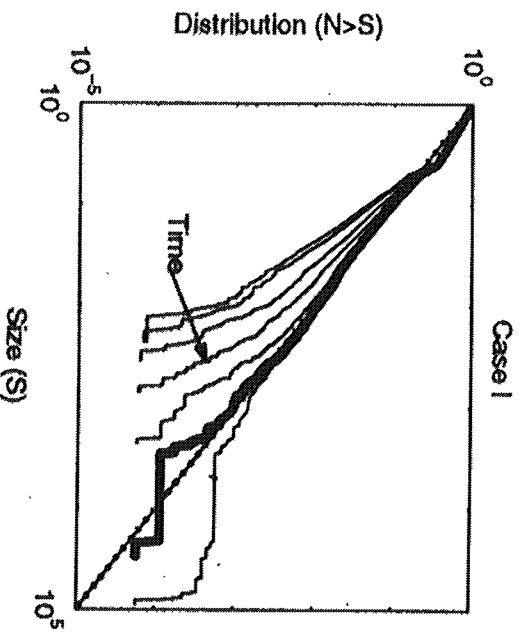
Timeline of Number of Events



Time (yr) $\times 10^3$

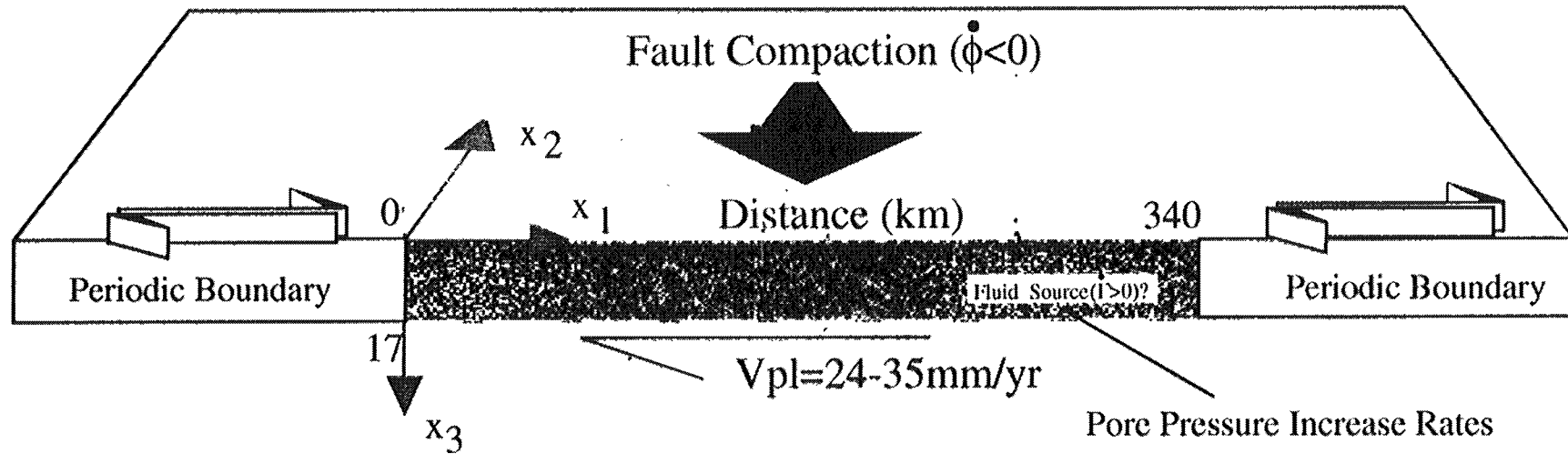
Miller & Nur (2000)
EPSL

CLUSTER SIZE DISTRIBUTION AT CRITICAL STATE



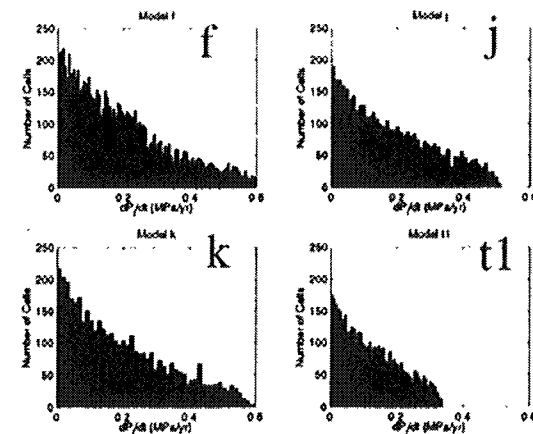
Willson & Nier (EPSL, 2000)

The Model



Model	V_{pl} (mm/yr)	μ_s	μ_d	$\Delta\sigma$ (%)
f	35	0.6	0.3	80
j	35	0.6	0.5	25
k	35	0.6	0.4	40
T1	24	0.6	0.5	25

Pore Pressure Increase Rates



MODEL ALGORITHM

Increase Shear stress

$$\tau_i = \frac{G}{2\pi} \sum_{j=1}^N k_{ij} (d_j - V_H t)$$

Increase Pore Pressure

$$\frac{dp}{dt} \Big|_{\text{no flow}} = \frac{(p - q)_i}{\phi_i \beta_i}$$

No

$$\tau > \tau_{\text{fail}} ?$$

YES

Redistribute Shear Stress

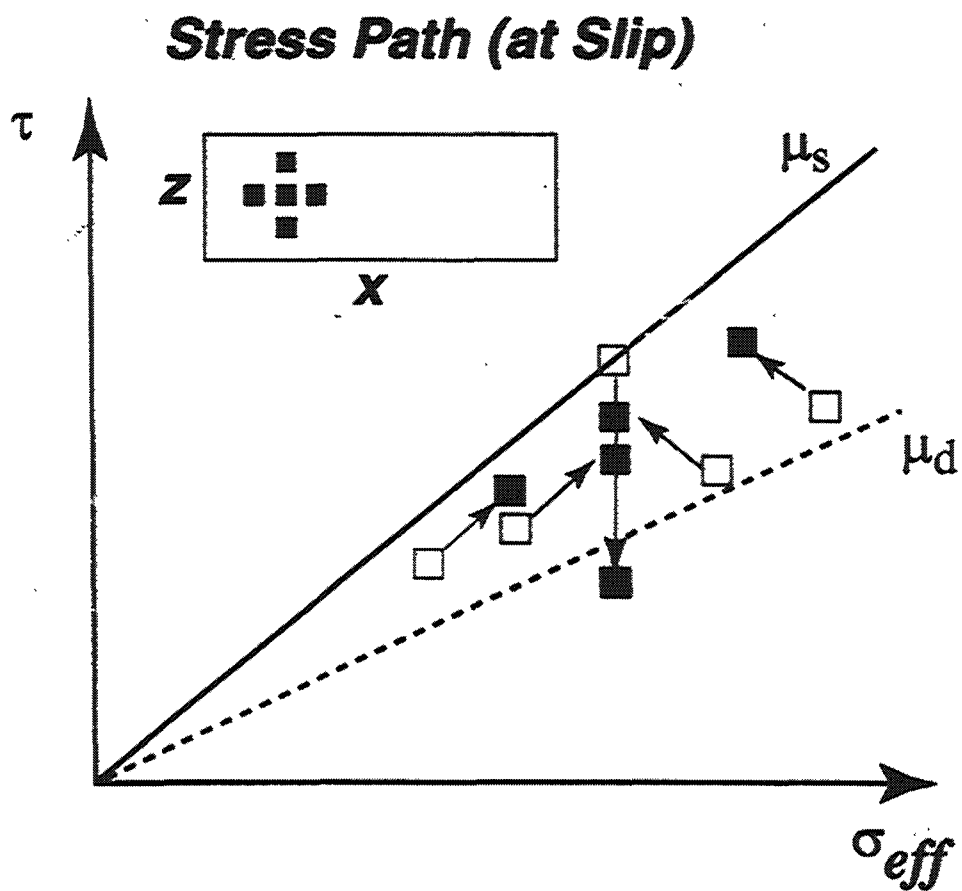
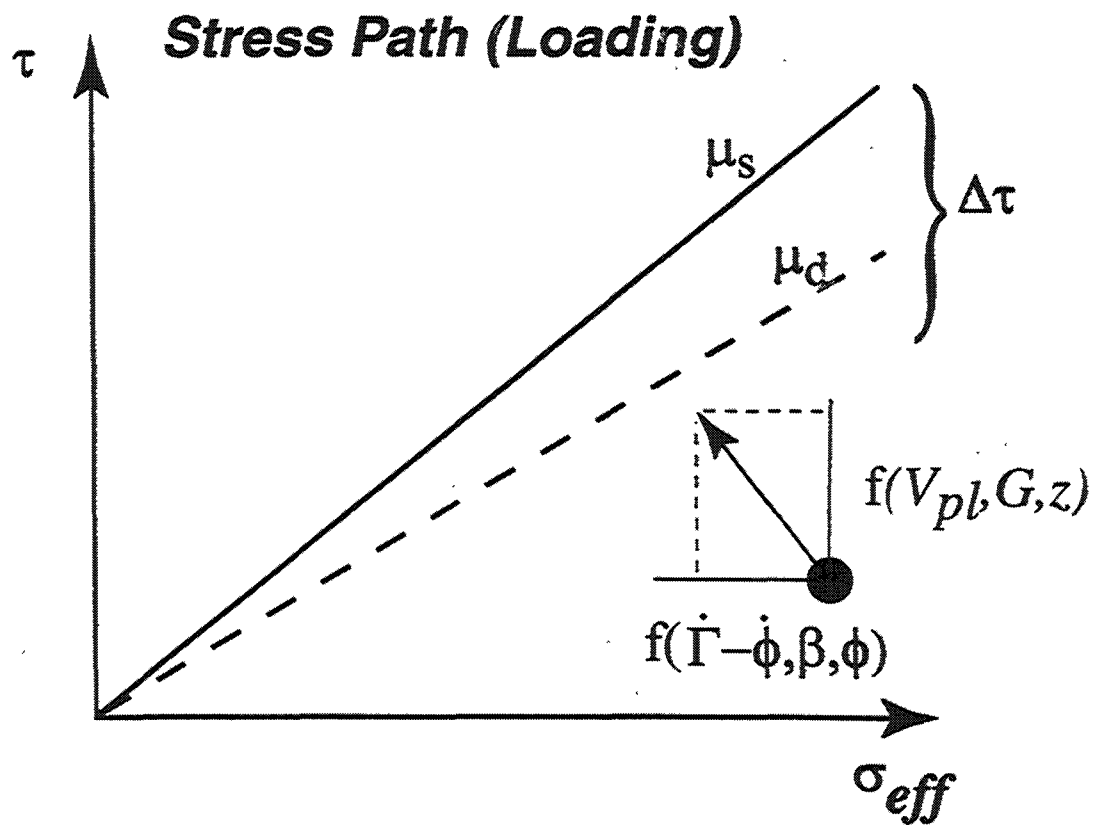
$$\tau_i = \frac{G}{2\pi} \sum_{j=1}^N k_{ij} d_j$$

Create Porosity

$$\frac{\partial \phi}{\partial t} = \frac{\beta_m (q - q_m) t}{\phi_c \Delta q_m}$$

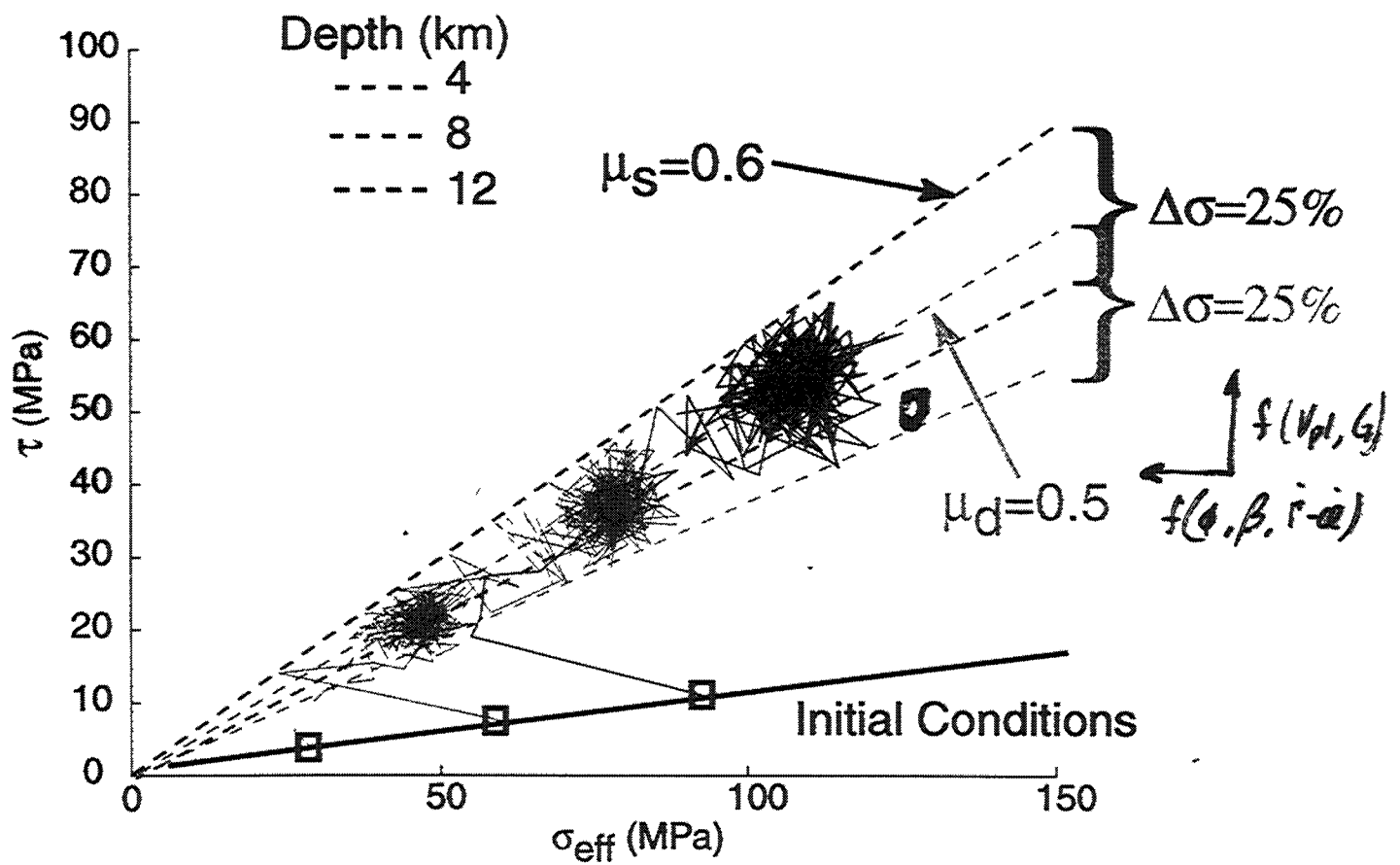
Redistribute Pore Pressure

$$\bar{p} = \frac{\sum_{i=1}^N (\phi_i \beta_i p_i - p_{g0})}{\sum_{i=1}^N (\phi_i \beta_i)} + p_{g0}$$

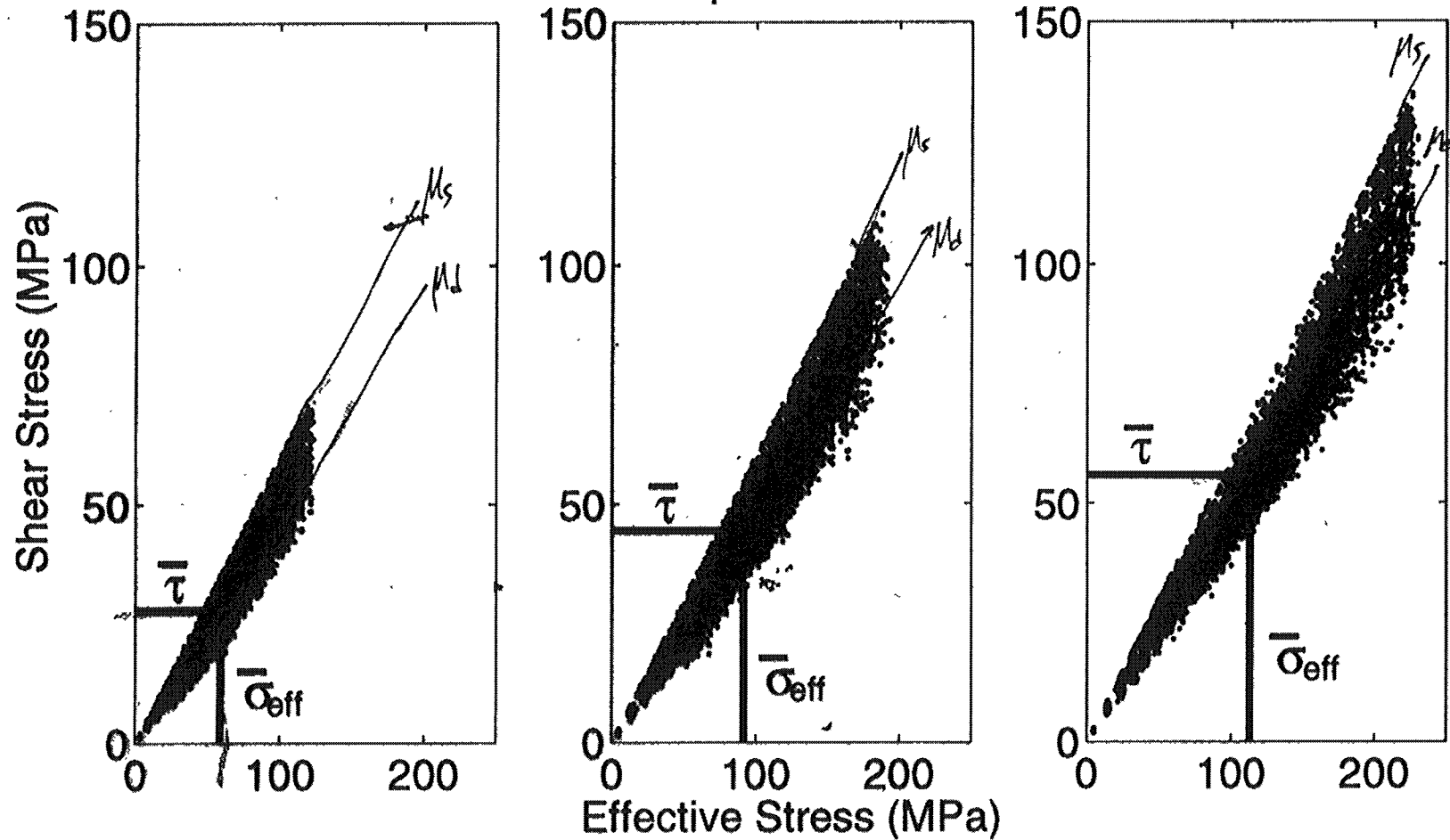


General Result: Model Finds Dynamic Equilibrium

$$f(V_p, \dot{\Gamma} - \dot{\phi}, \Delta\sigma, \mu)$$



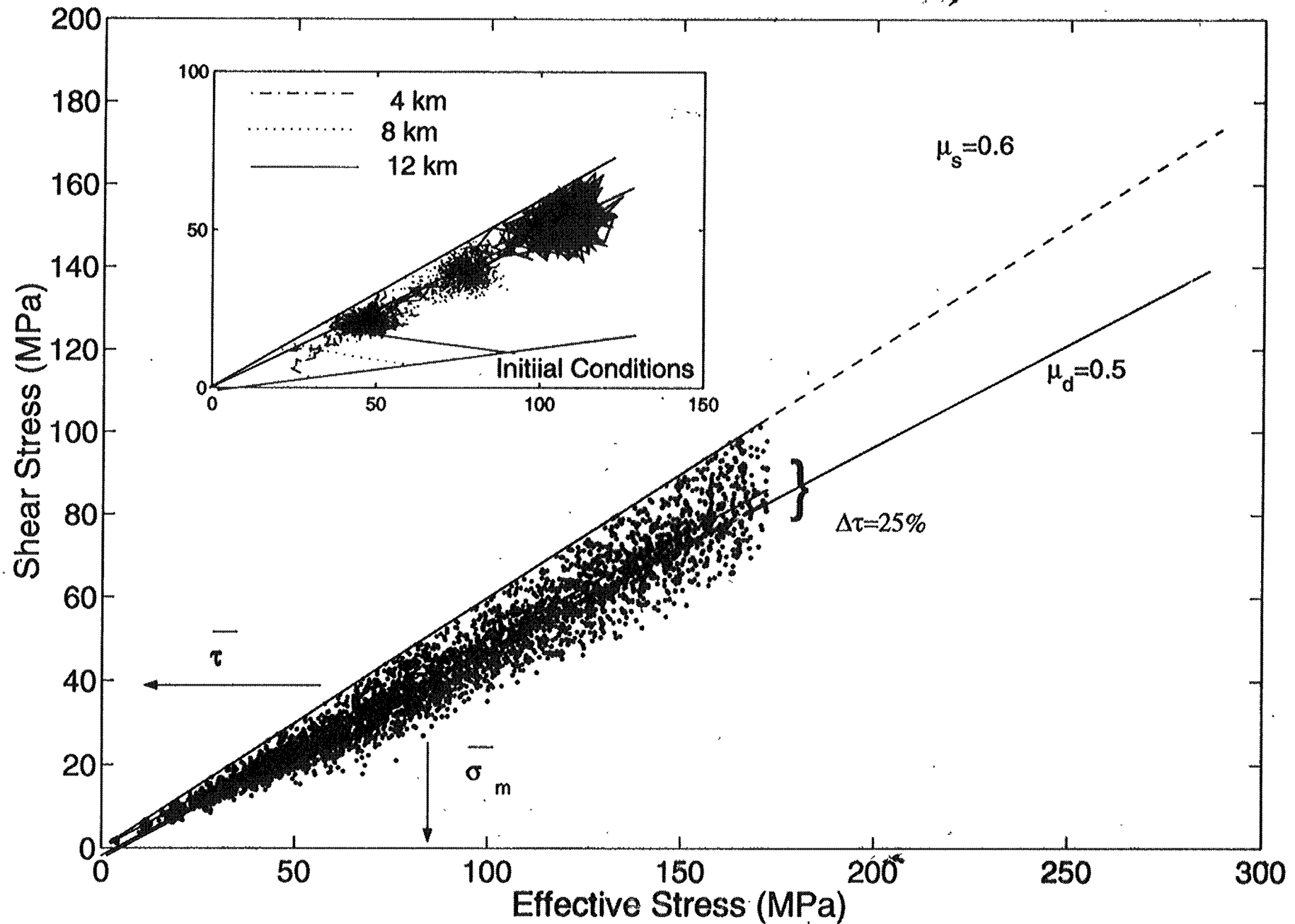
Evolved "Equilibrium" Stress State



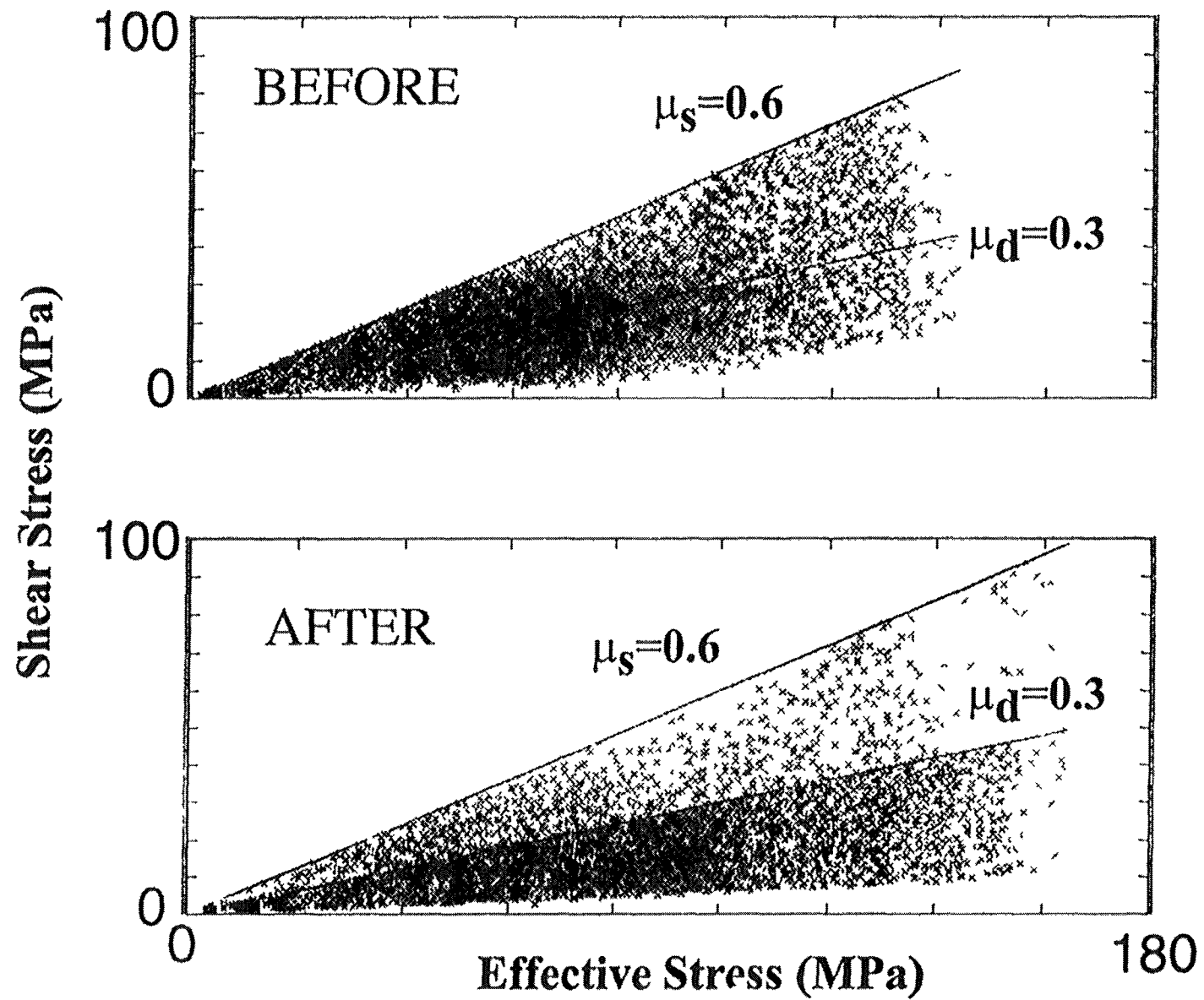
$$\Delta \bar{\tau}_d = (C_1 \mu_s - \mu_d) \bar{\sigma}_e = (C_1 \mu_s - \mu_d) \rho g \bar{z} (1 - \lambda)$$

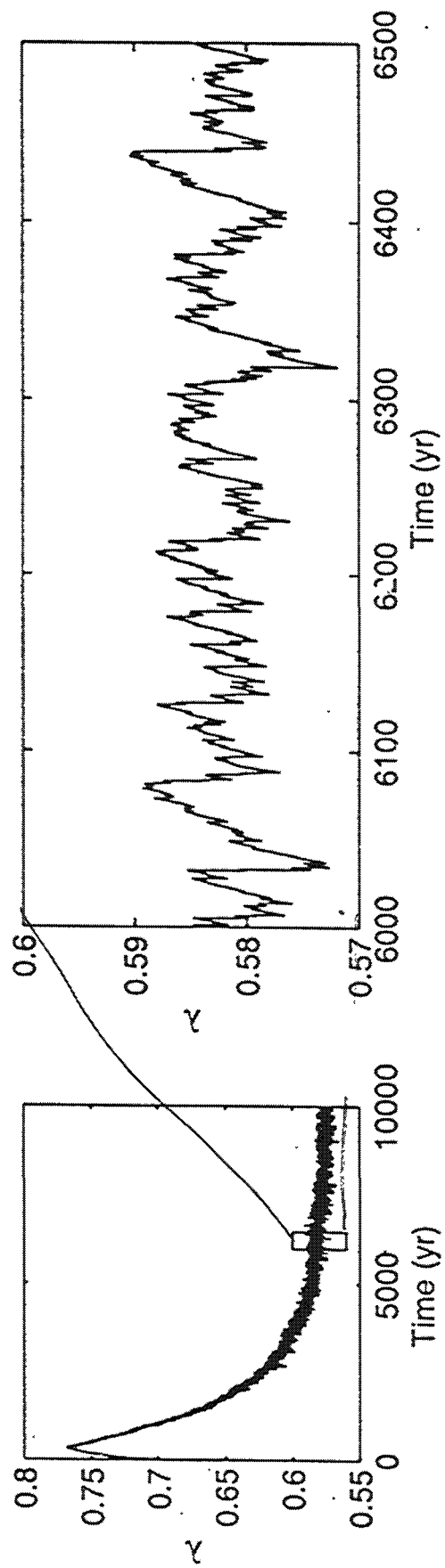
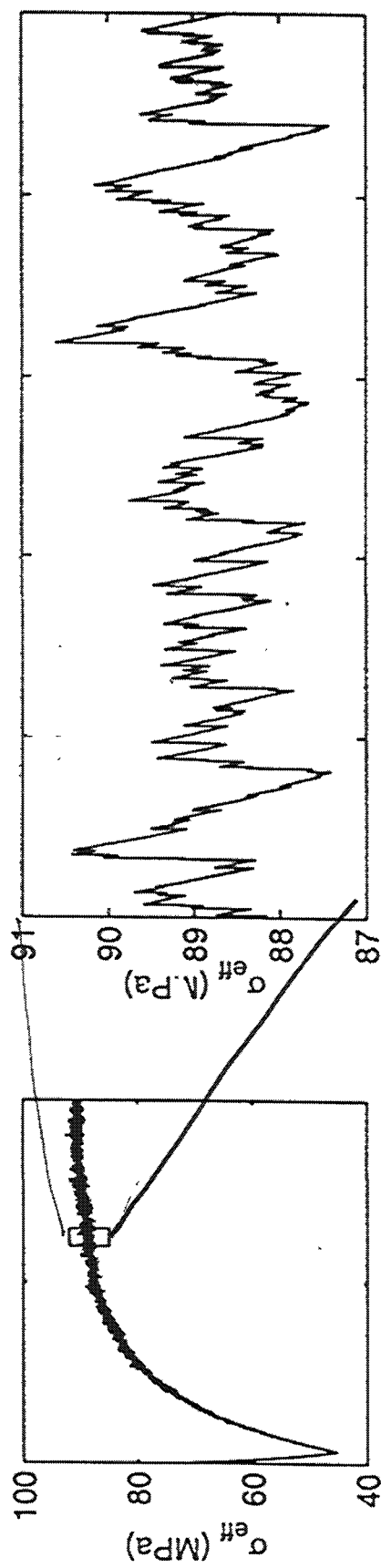
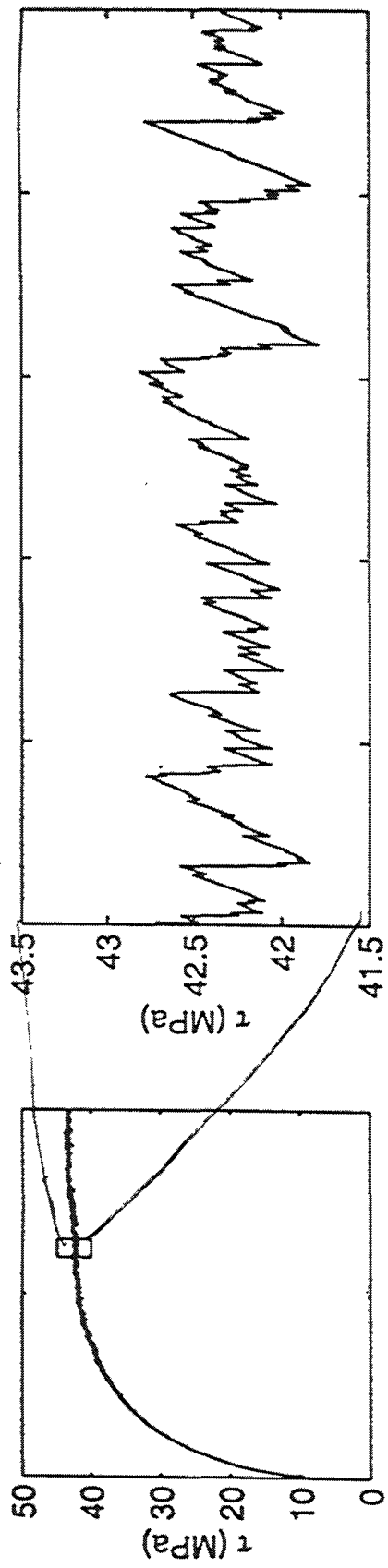
$$\bar{z} \rightarrow \frac{1}{2} W \Rightarrow \Delta \bar{\tau}_d = \frac{1}{2} (C_1 \mu_s - \mu_d) \rho g W (1 - \lambda)$$

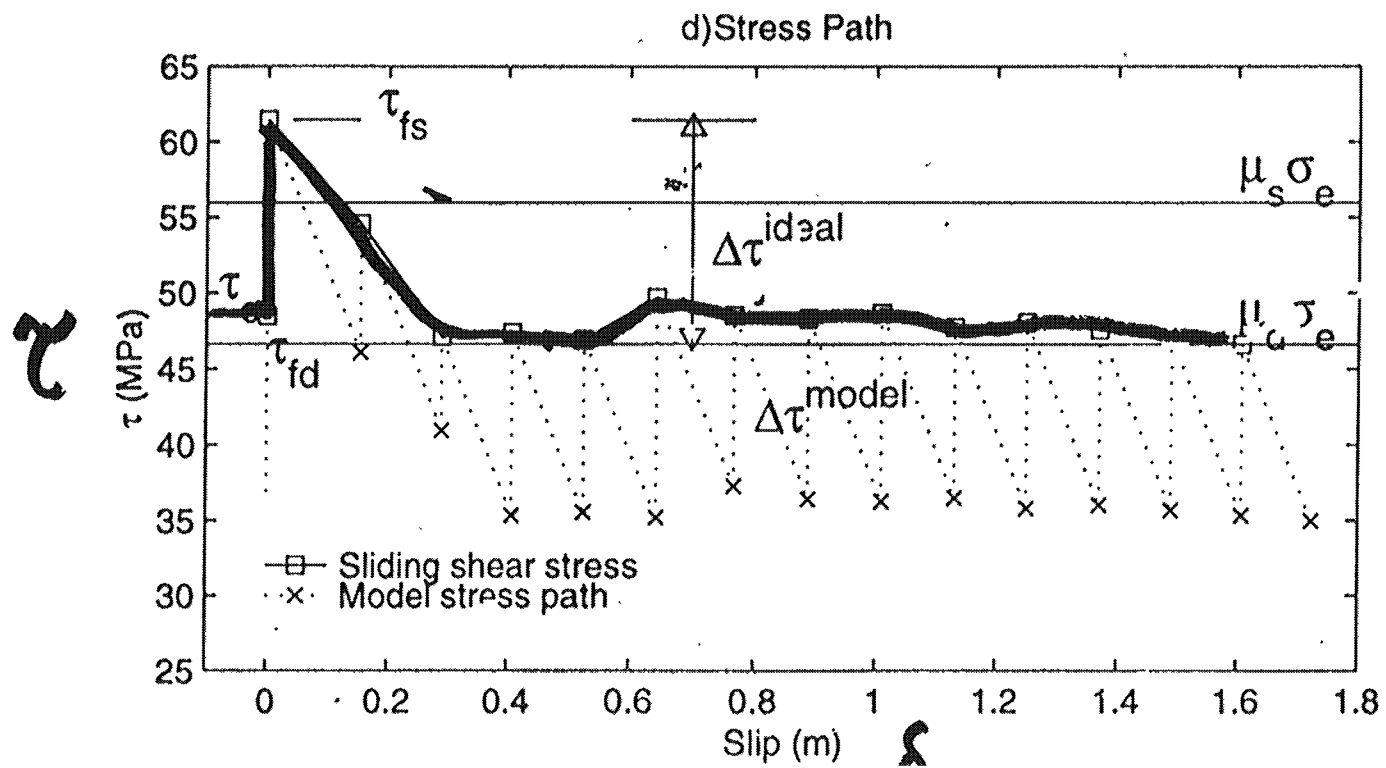
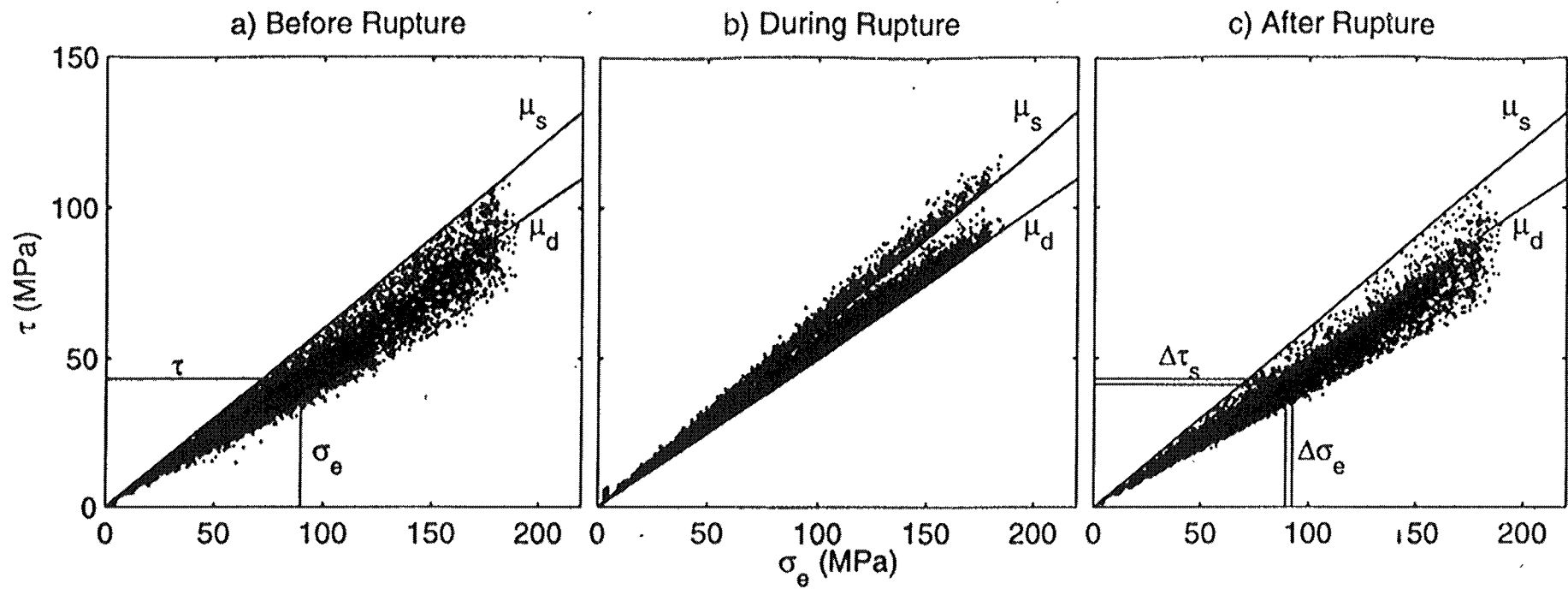
Stress Paths & Dynamic Equilibrium



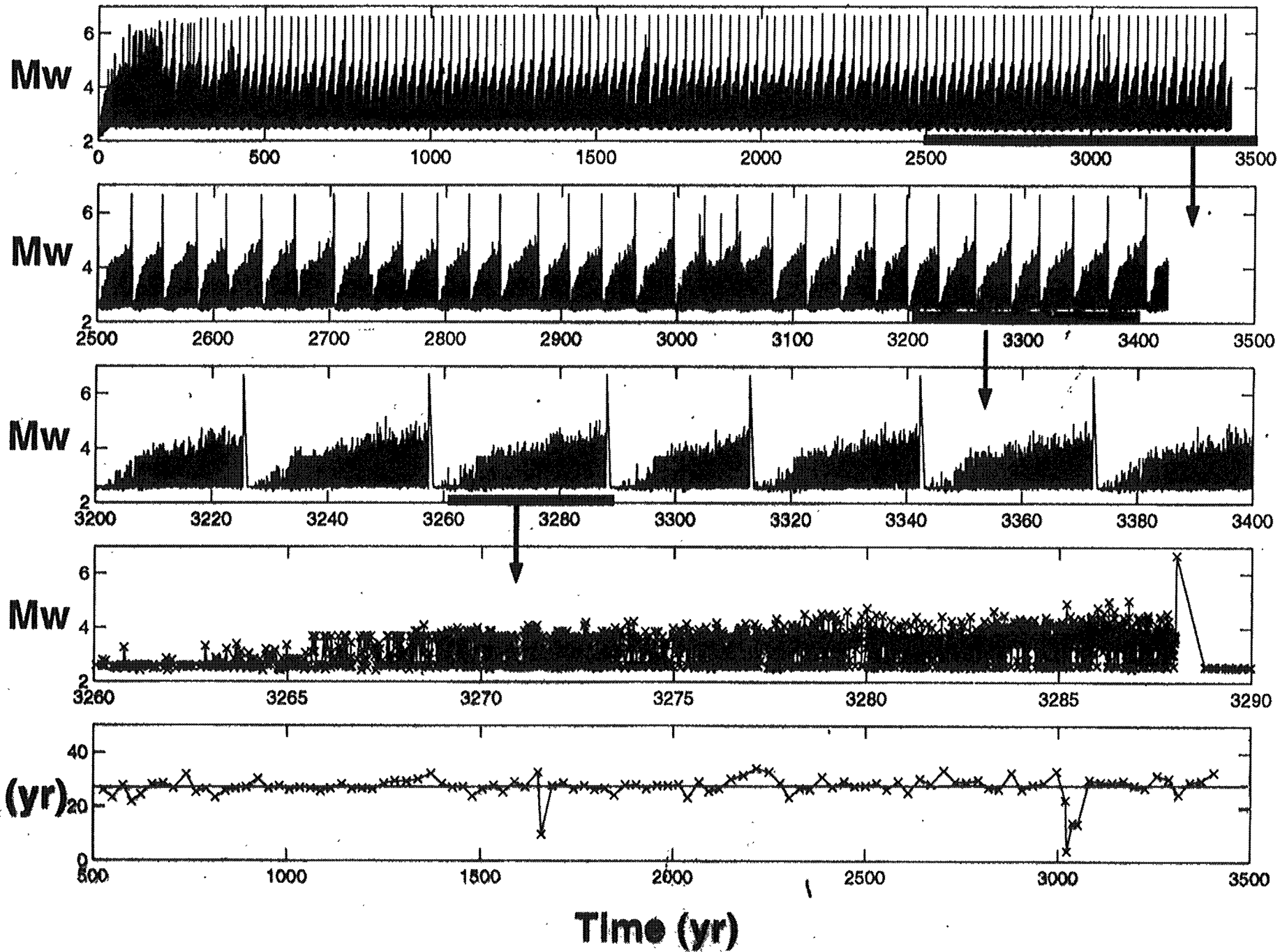
Stress-state before and after a large rupture

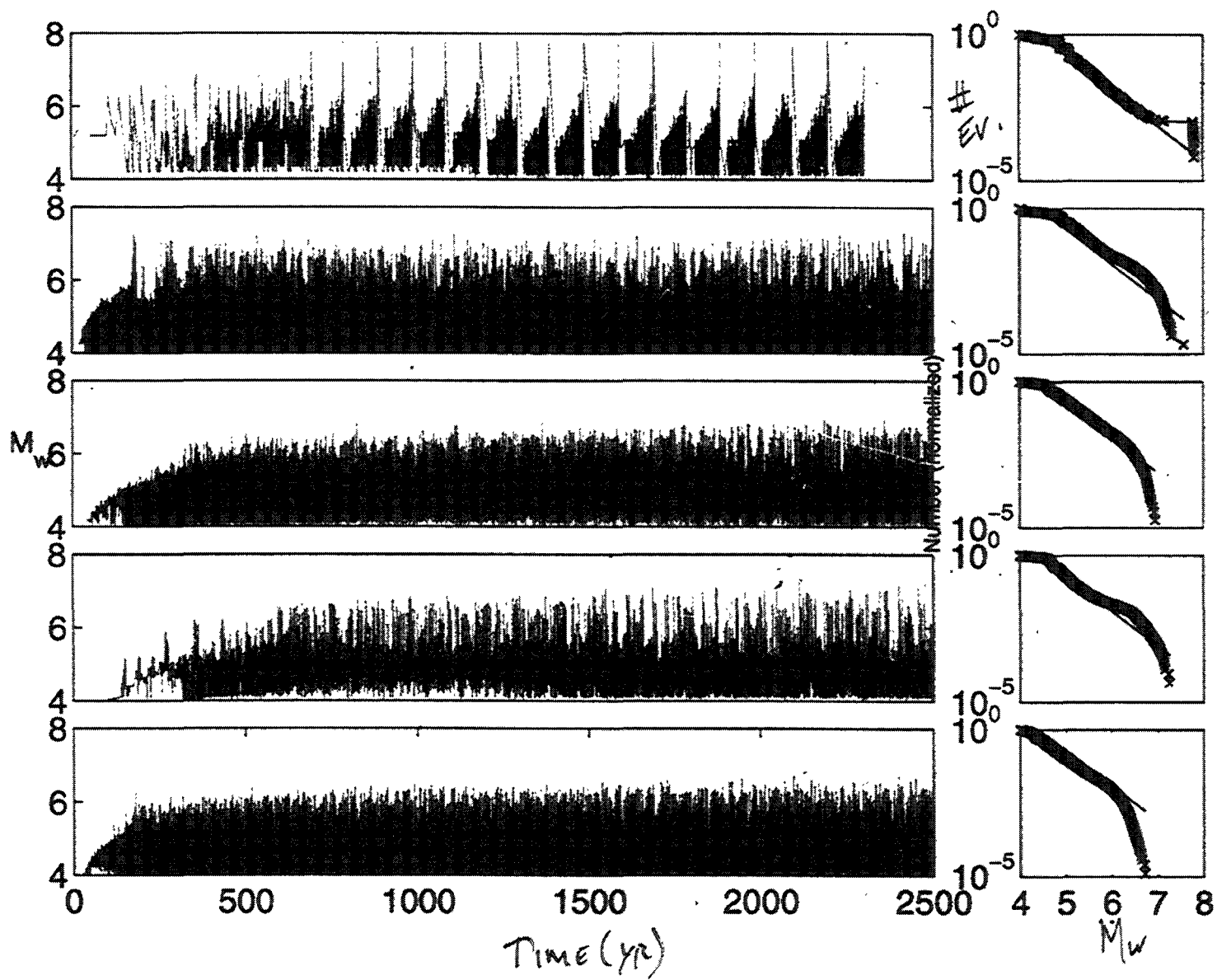


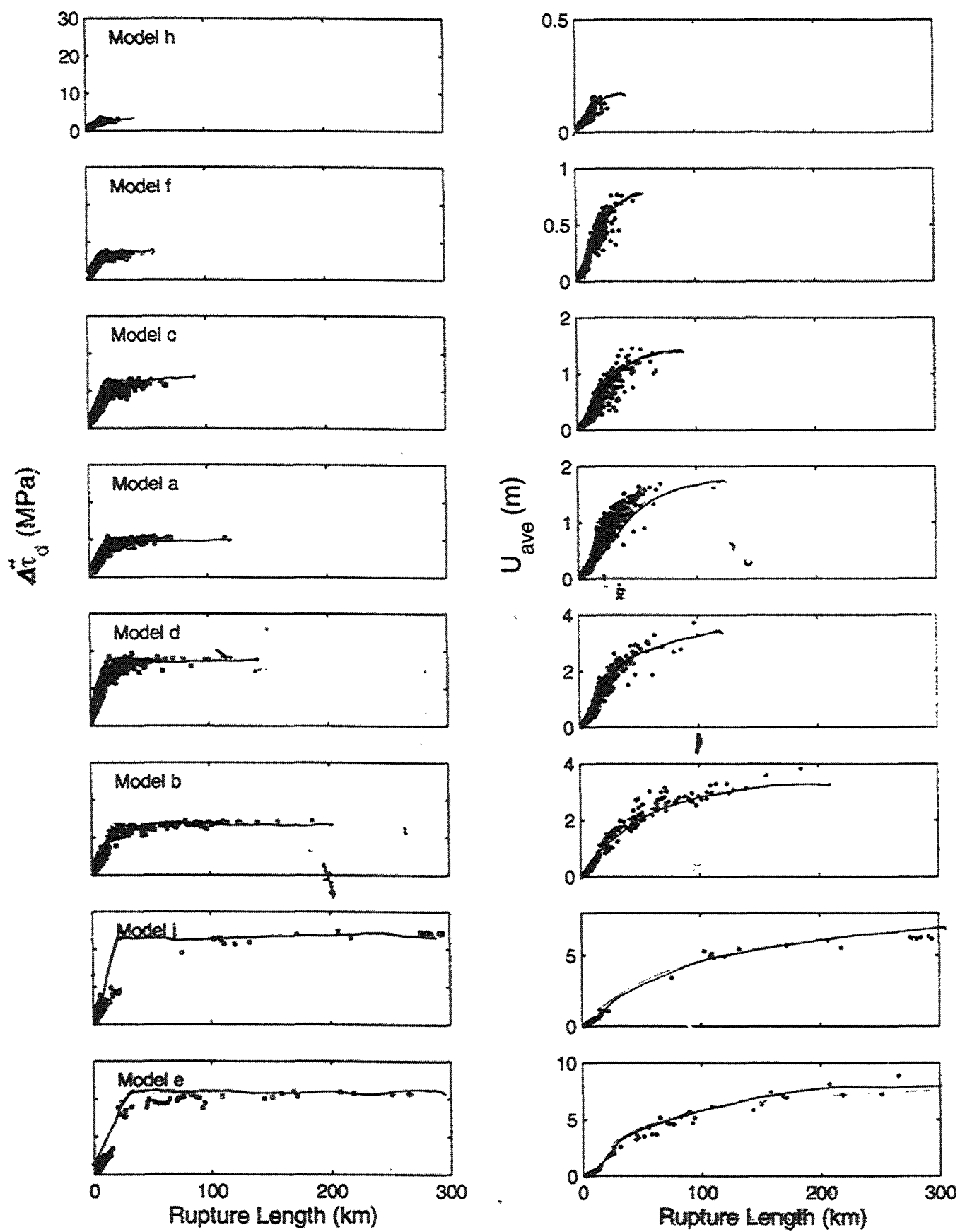




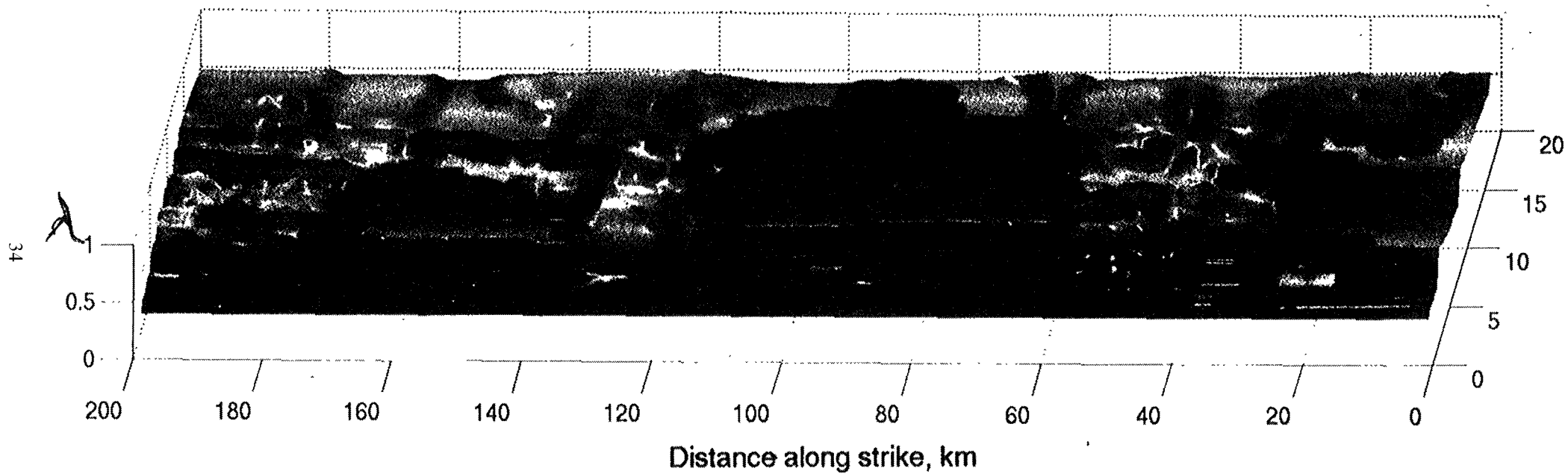
Repeat Times and the Earthquake Cycle

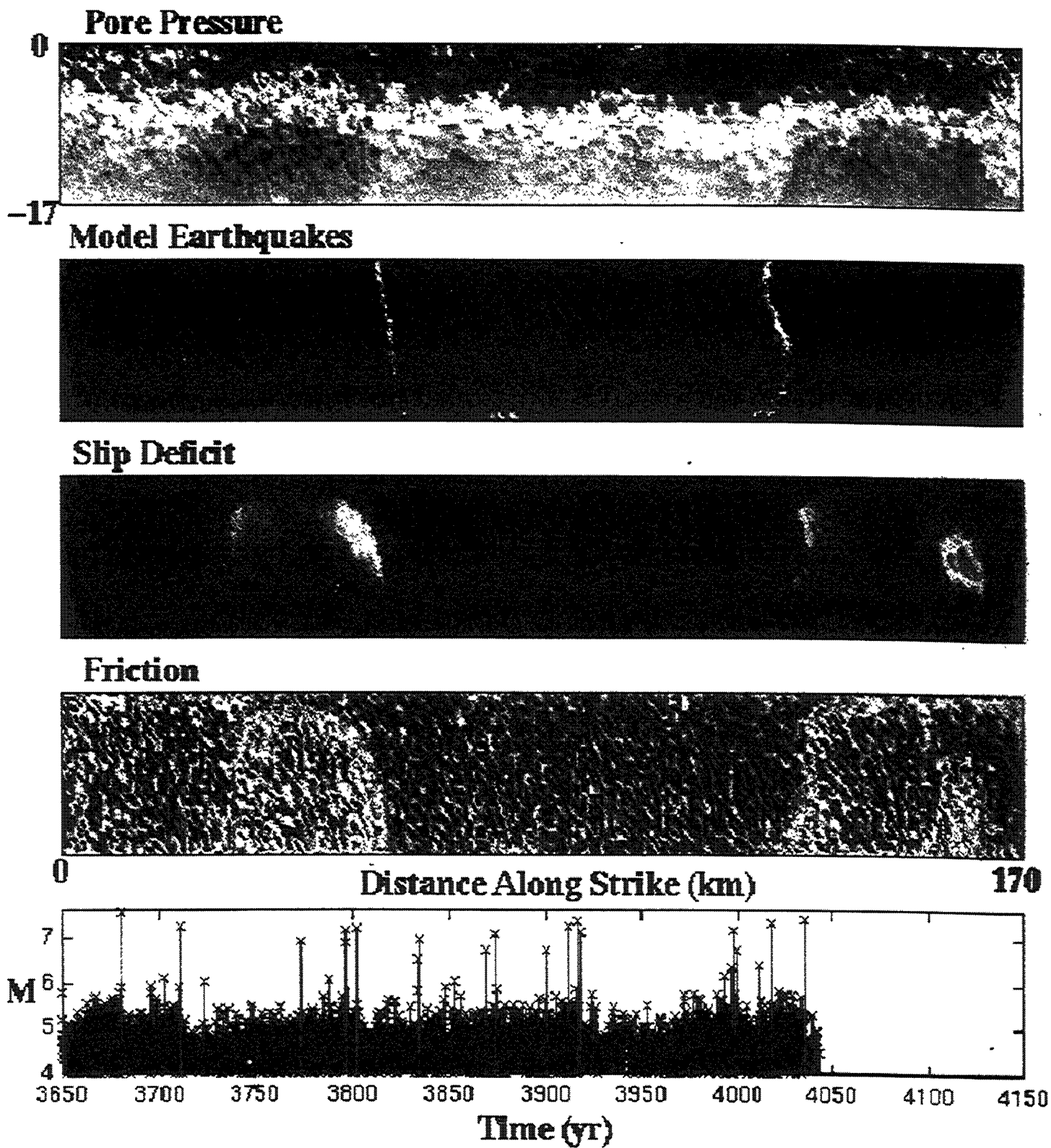


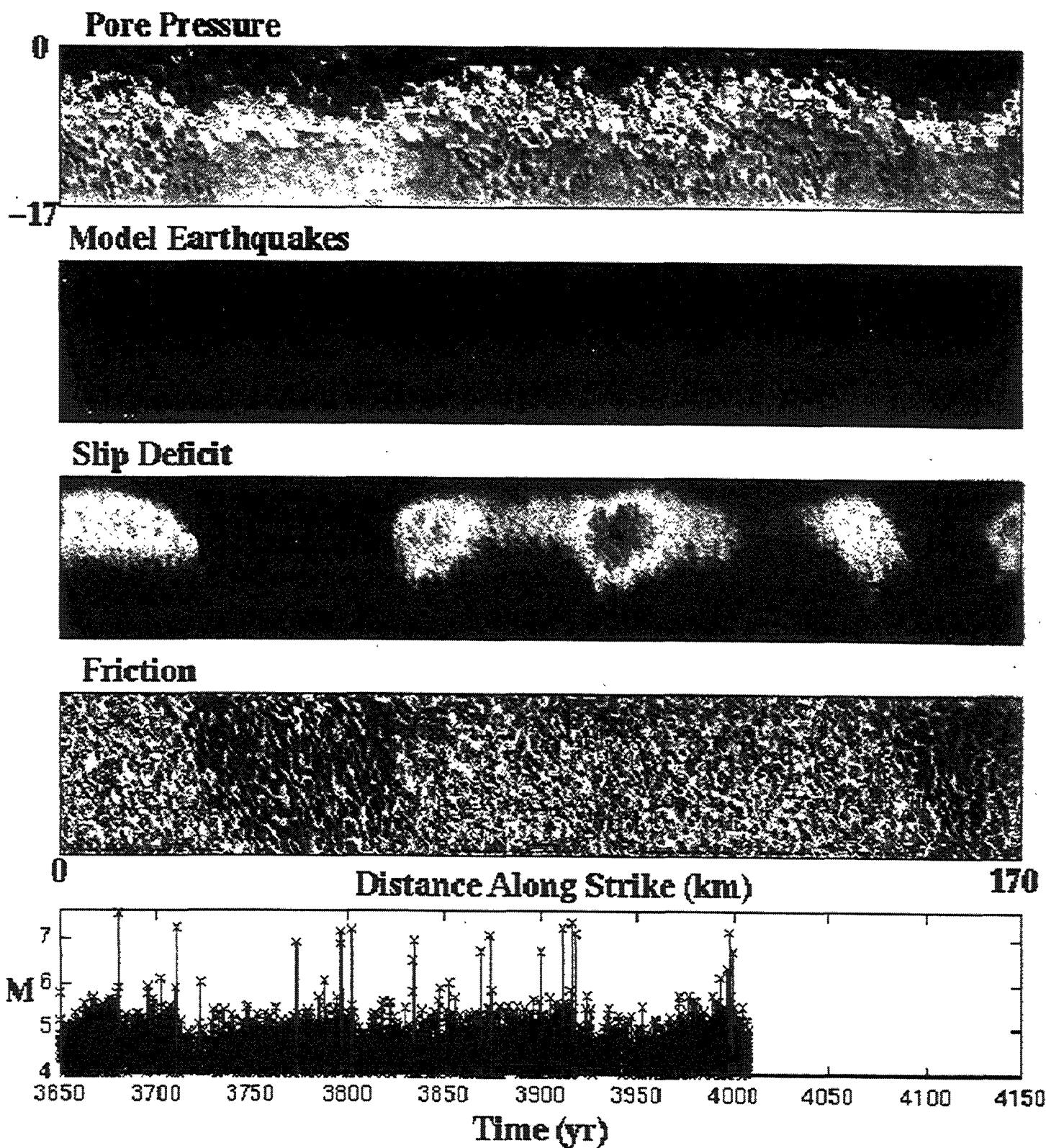




PORE PRESSURE ALONG FAULT







Comparison of Moment-Rupture Area with Historical Strike-Slip Catalog (Wells and Coppersmith, '94)

