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H4.SMR/1330-29

" Sixth Workshop on Non-Linear Dynamics and
Earthquake Prediction"

15 - 27 October 2001

**Earthquake Scaling and the Strength
of Seismogenic Faults**

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Lecture I: The dynamical influence of fluids on seismicity and faulting

- Rock and fluid mechanics
- Dislocations and brittle faulting
- Stress transfer models
- The hydraulics of fault zones
- Coupling fluid flow to large scale tectonics
- Earthquakes as a coupled shear stress, high pore pressure dynamical system
- The properties of large model earthquakes

Lecture II: Earthquake scaling and the strength of seismogenic faults

- Earthquake scaling: Small earthquakes
- Earthquake scaling: Large earthquakes
- Pore pressure as a fundamental scaling parameter
- Application of result to global earthquake catalogs
- Can pore pressures be inferred from earthquake rupture properties

Lecture III: Neotectonics and earthquake generation along complex fault systems

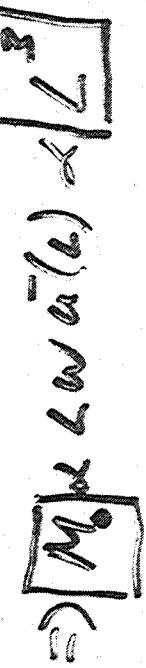
- The North Anatolian Fault Zone
- The San Andreas Fault System
- Building a forward model of the earthquake process
- Fault Interaction

Earthquake scaling

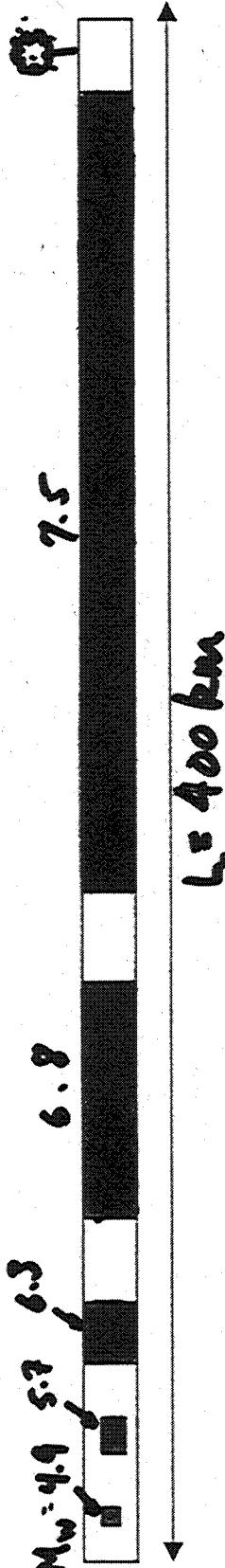
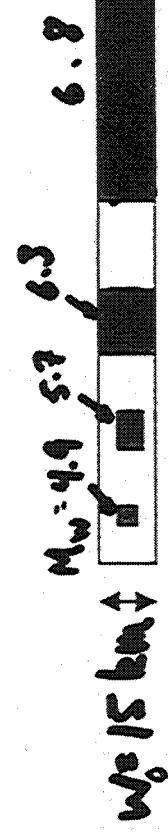
$$M_0 = \mu L \omega^2$$

μ : shear modulus
 L : rupture length

ω : constant
 $\Rightarrow \bar{\omega}(L)$



μ : shear modulus
 L : rupture length
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 $\Rightarrow \bar{\omega}(L)$
 $\Delta \varepsilon = C_m \left(\frac{\omega}{L} \right)$
 λ : scale length



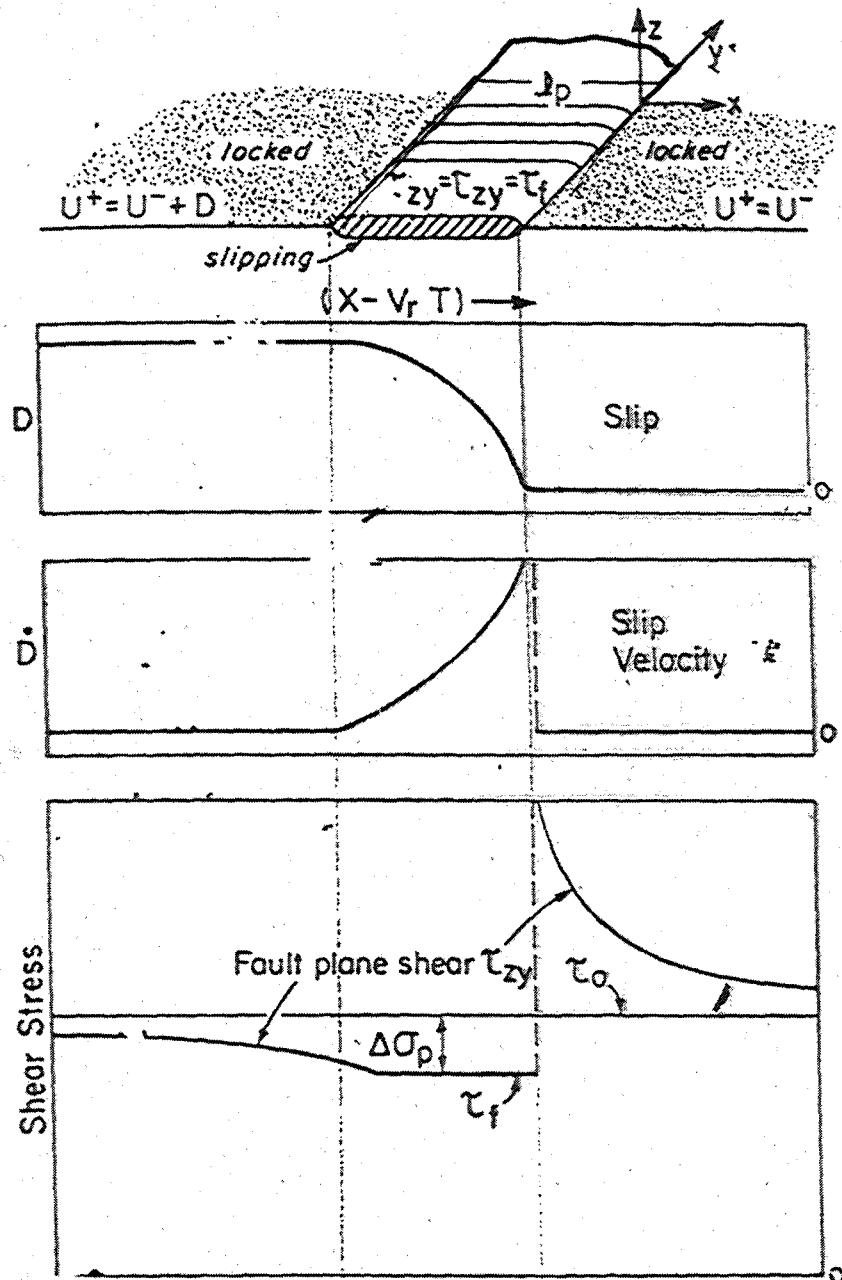
$L A R S E$: $\frac{L}{w_0} > 1$ if $\bar{\omega}(L) \Rightarrow M_0 \propto L \ln \bar{\omega}(L) \propto L^2 \rightarrow L\text{-model}$

$w = \text{const.}$

$\Delta T \propto \text{constant}$

if $\bar{\omega} \propto \bar{\omega}(L)$ $\Rightarrow M_0 \propto L \ln \bar{\omega} \propto L \rightarrow \omega\text{-model}$

SELF-HEALING PULSES OF SLIP IN EARTHQUAKE RUPTURE



Healer layer

