

Advanced Course:
CLIMATE CHANGE IN THE MEDITERRANEAN REGION
PART I: PHYSICAL ASPECTS
(12 - 16 March 2001)

**"Effects on Climate of Vegetation, including
Deforestation and Afforestation"**

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These are preliminary lecture notes, intended only for distribution to participants

Effects of vegetation on climate

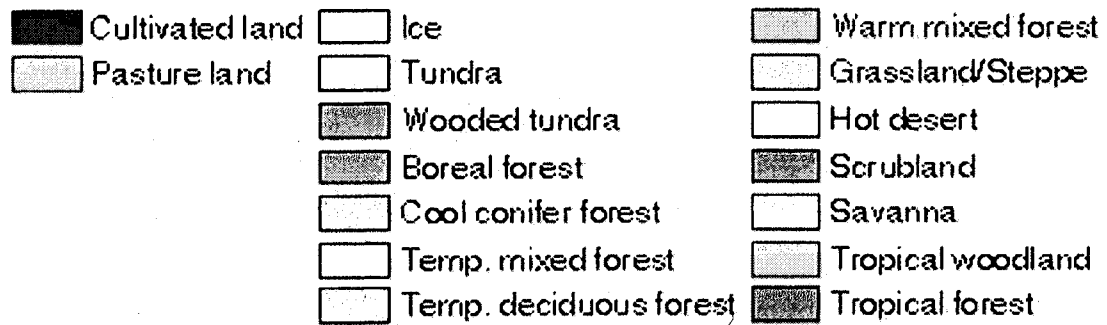
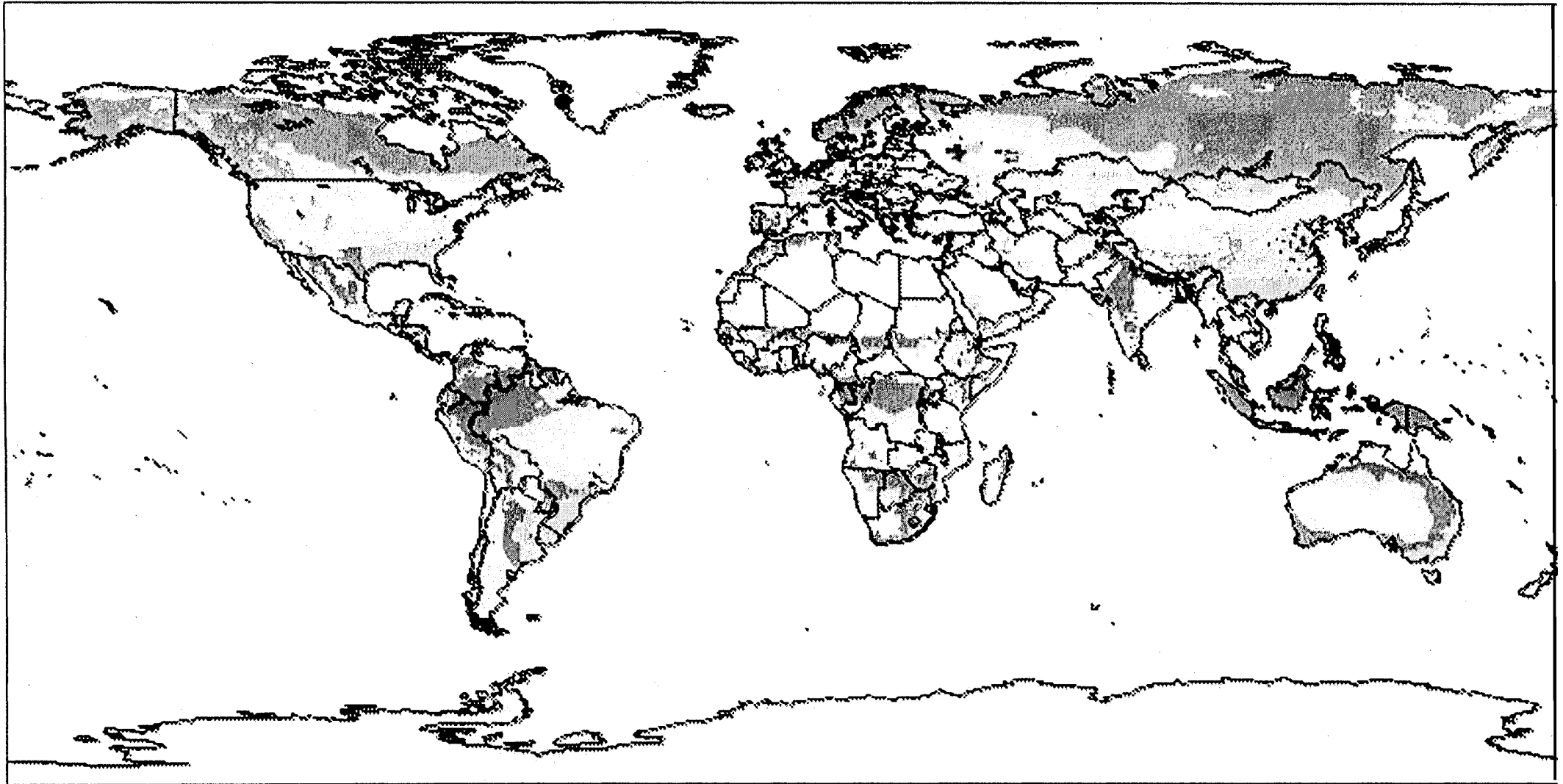
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Effects of climate on vegetation

Overview

- current vegetation distribution
- vegetation as a function of climate
- anthropogenic land use change
- description of vegetation effects in GCMs
 - energy and water cycle
- sensitivity of energy and water cycle to deforestation
- vegetation and the carbon cycle
- coupled climate-carbon models
 - first results

year 1700



Vegetation is part of the climate system

Vegetation is part of the carbon cycle,
and the water cycle,
and has radiative effects

changes in vegetation may change
the water cycle,
the energy budget,
the carbon cycle

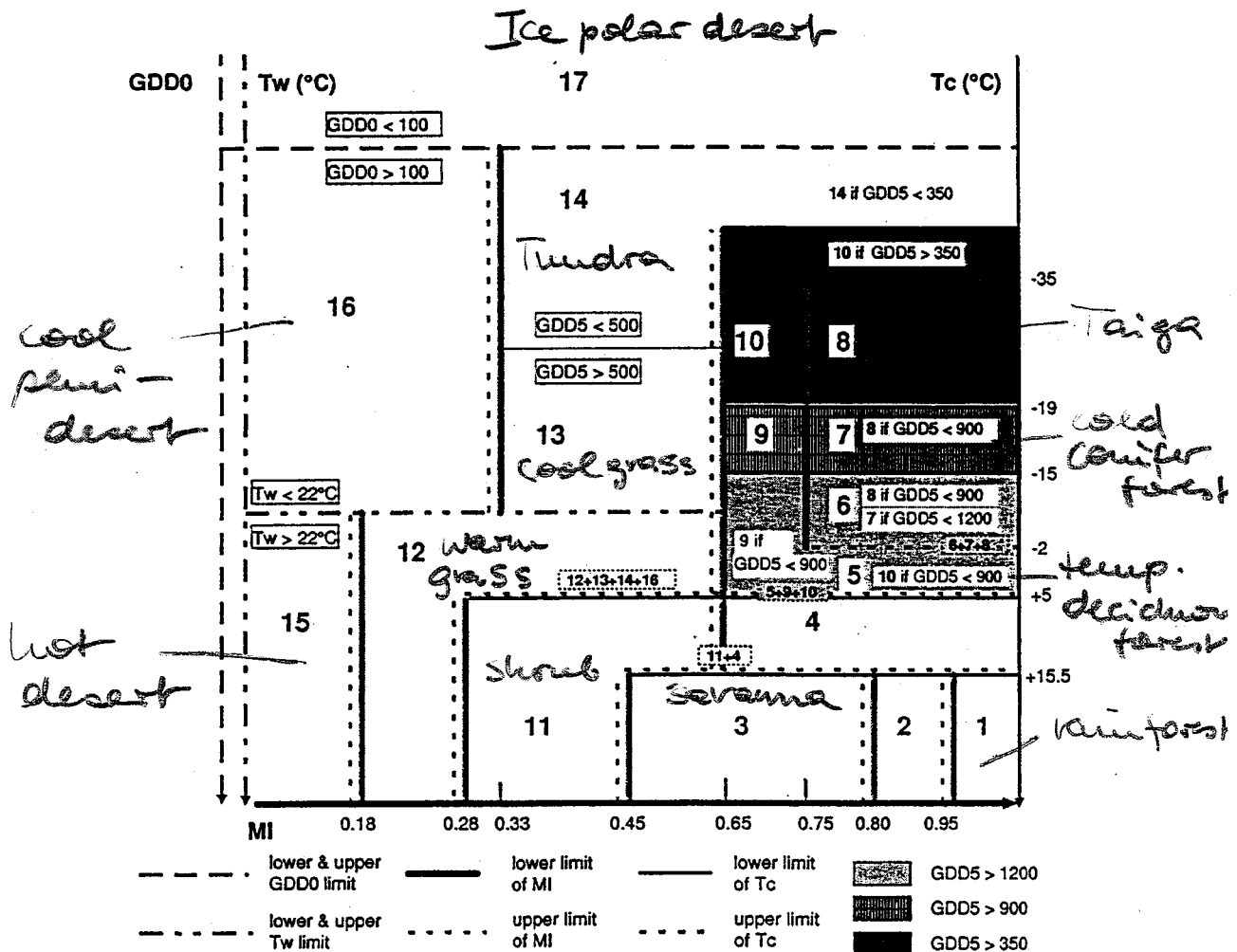
changes in the carbon cycle are
suspected to change
the climate.

changes in the climate may
change the water cycle,
change the energy budget,
change the vegetation

changes in the vegetation may
change the carbon
cycle (terrestrial sink)

BIOME model.

static



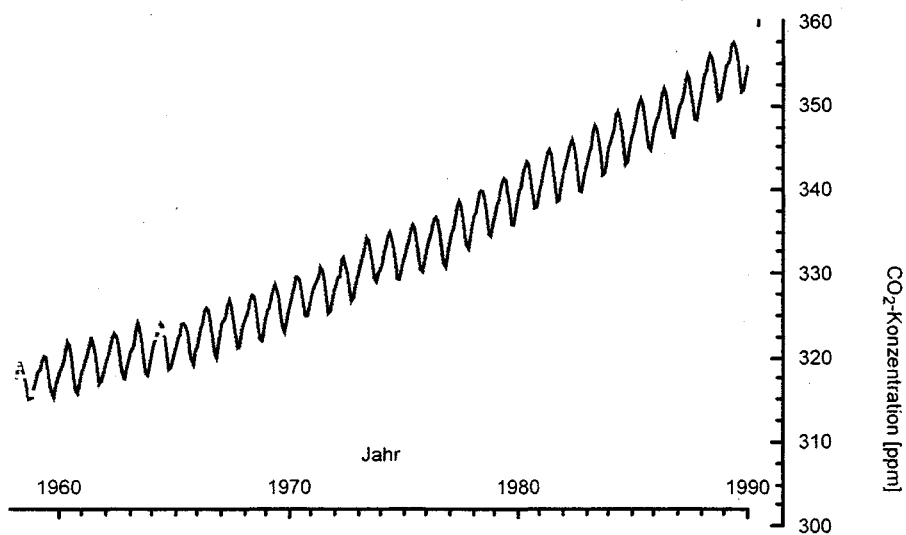
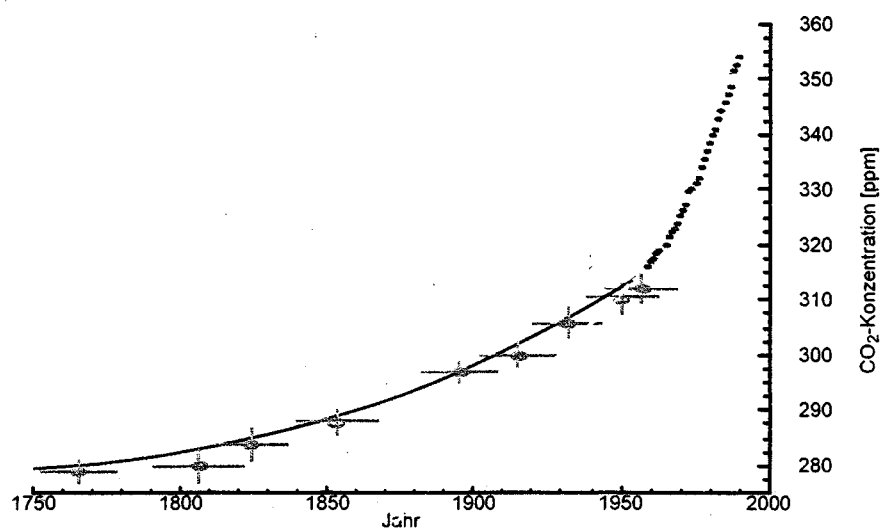
Colin Prentice, Jma

Figure 16.1 Environmental constraints imposed to 17 biomes (bold numbers) in the BIOME model (see Table 16.2 for the list of biomes). Each biome in the diagram has one or more environmental limits: moisture index MI , temperature of the coldest month T_c , growing degree day base 5°C and 0°C , $GDD5$ and $GDD0$, respectively, and temperature of the warmest month T_w . The biomes are limited by lower and/or upper environmental limits, but there is not necessarily one lower and one upper environmental limit allocated to each biome. The dotted boxes with bold numbers represent the biomes for which there is an upper T_c limit (e.g. the upper T_c limit of biomes 6, 7 and 8 is -2°C).

The primary input variables are monthly precipitation, surface temperature, the total amount of cloud and soil data. In this study the soil water capacity is set to 15 cm and so no additional soil data are employed. The Holdridge model was tuned (Figure 16.2) in order to match the distribution of vegetation given by the BIOME model when observed climate data are used. This tuning was done using a high resolution dataset of climate (i.e. 0.5° by 0.5°).



Changes in CO₂ Concentration



Vegetation

wind field :

roughness length

radiation :

albedo (snow)

water cycle :

vegetation fraction
vs. bare soil

vegetation types

by LAI

by minimum
stomatal

resistance

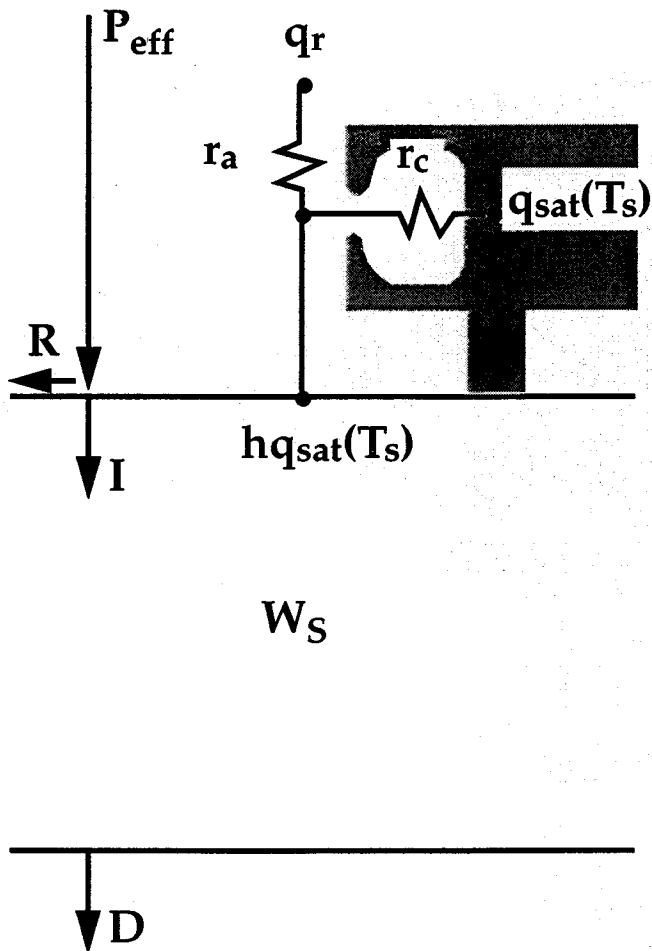
by root depth

Soil character. :

Soil organic matter

Carbon cycle

Land Surface Moisture Fluxes in the ECHAM GCM



Symbols:

P_{eff} : effective precipitation

R : surface runoff

I : infiltration

q_r : spec. humidity at reference height

q_{sat} : saturation value of q

T_s : surface temperature

W_s : soil moisture

D : drainage

h : relative humidity

Resistances:

r_a : aerodynamic

r_c : bulk stomatal

Bucket hydrology: One soil moisture layer of depth W_{Smax}

$$\rho_w \frac{\partial W_s}{\partial t} = P_{eff} - E_{eff} - R - D + M_{eff}$$

E_{eff} : effective evapotranspiration

M : snow melt

ρ_w : density of water

$$E_{eff} = E_{bs} + E_{tr}$$

horizontal distribution of land surface characteristics

Bare soil evaporation:

$$E_{bs} = \rho_w \frac{q_r - h q_{sat}}{r_a}$$

with

$$h = \max \left[\frac{1}{2} \left(1 - \cos \frac{\pi \cdot W_s}{W_{Smax}} \right), \min \left(1, \frac{q_r}{q_{sat}} \right) \right]$$

Transpiration:

$$E_{tr} = \rho_w \frac{q_r - q_{sat}}{r_a + r_c}$$

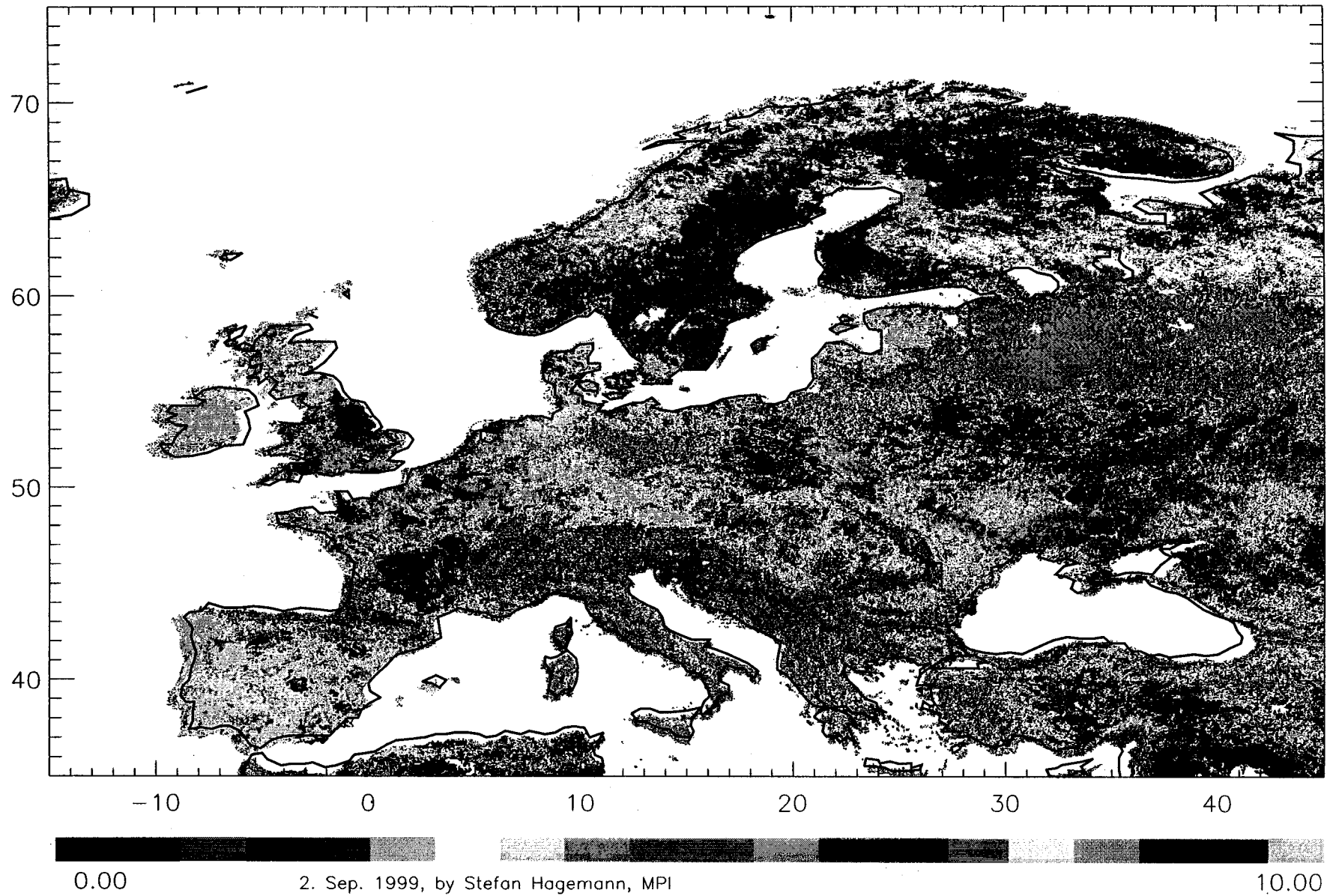
with

$$r_c = f(PAR, W_s)$$

PAR: photosynthetically active radiation

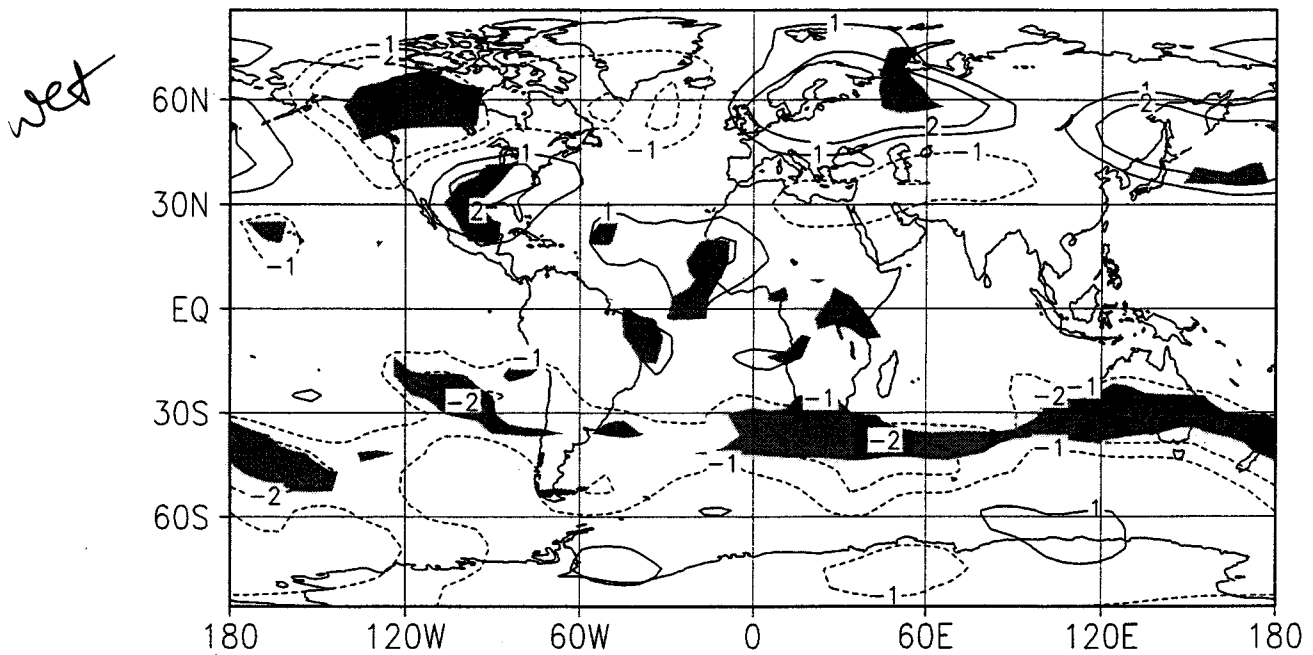
IGBP - Dis

LAI Growing Season at 0.05 degree resolution (error in IDL LSM mask)

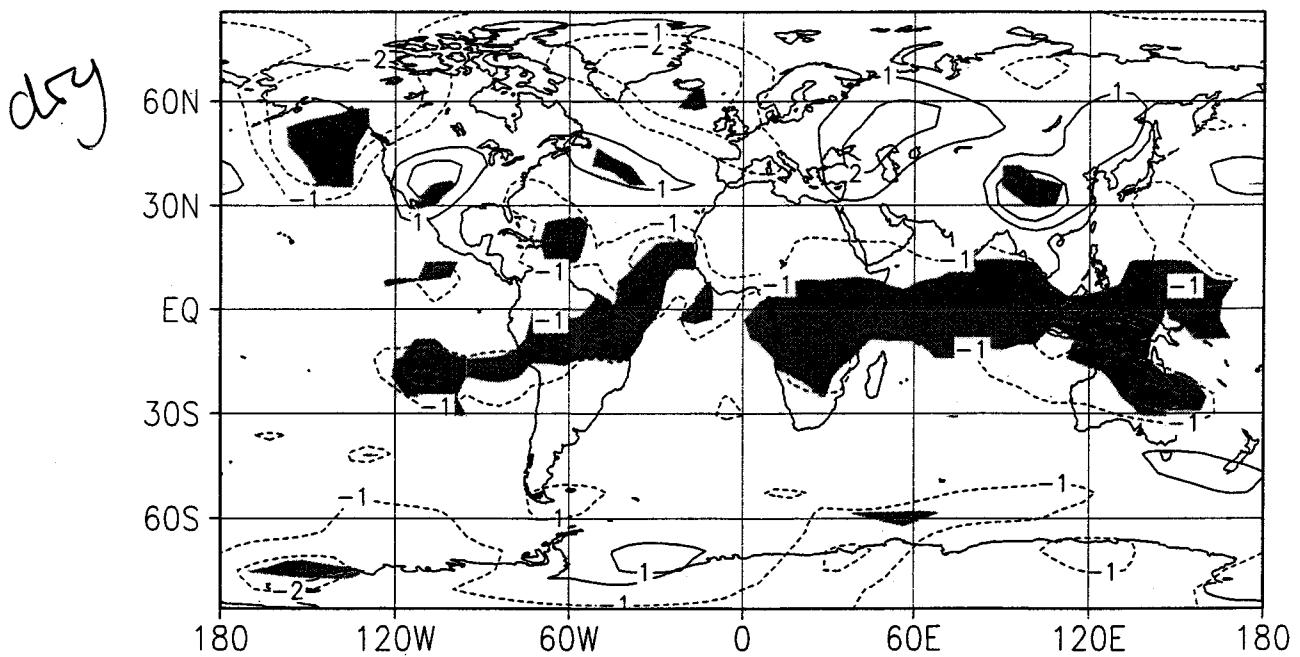


temperature [°C] 500hPa Difference January 02–05

23008–23005



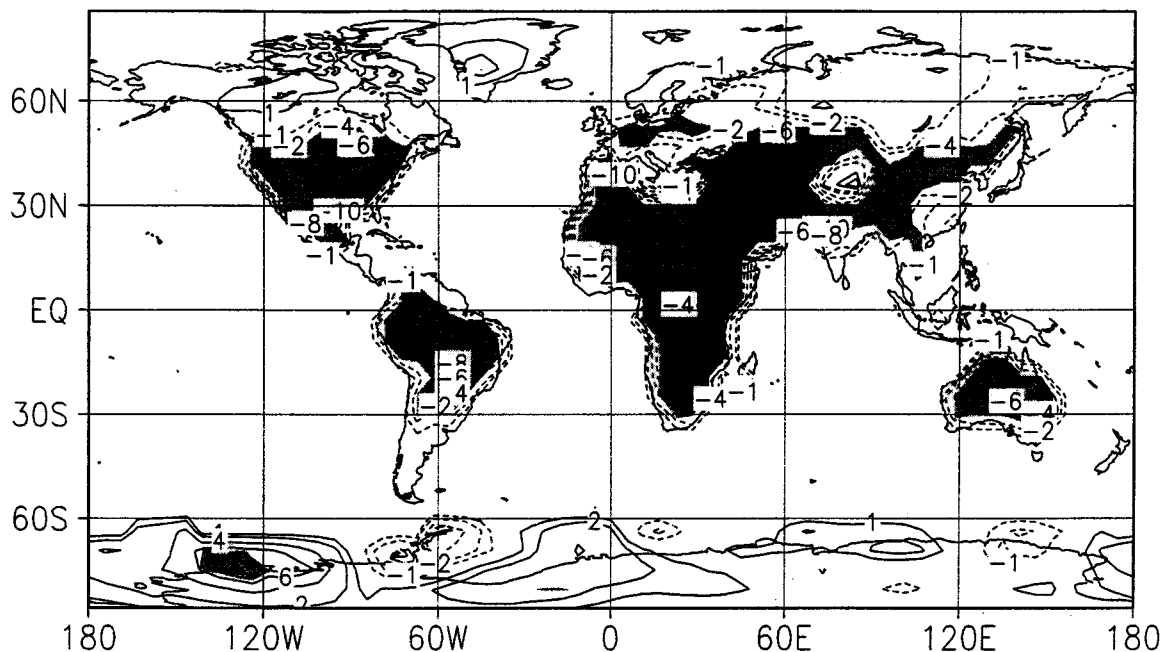
23009–23005



surface temperature [°C] Difference July 02–05

23008–23005

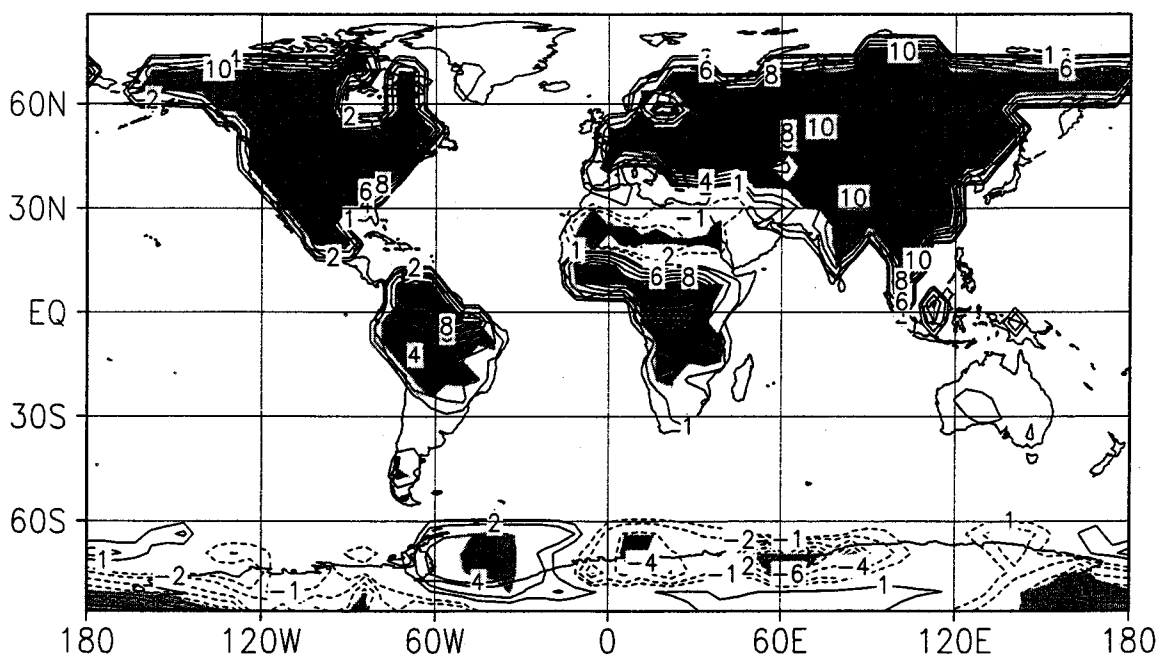
wet



97% significance level

23009–23005

dry



Deforestation / Afforestation

a) a land use change
→ carbon cycle

b) a change of atmospheric
boundary conditions

effects on energy budget

on water cycle

on carbon cycle

examples: Amazon, Indonesia, Africa
Sahel
Mediterranean

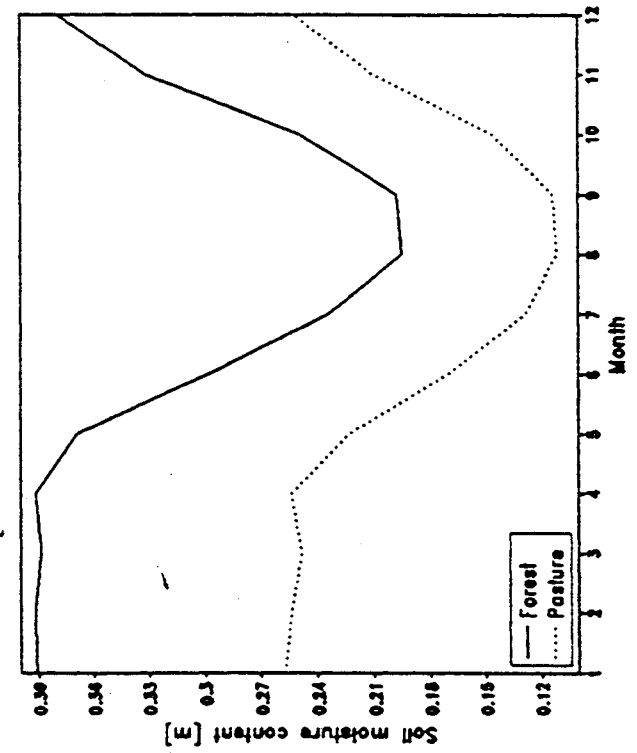
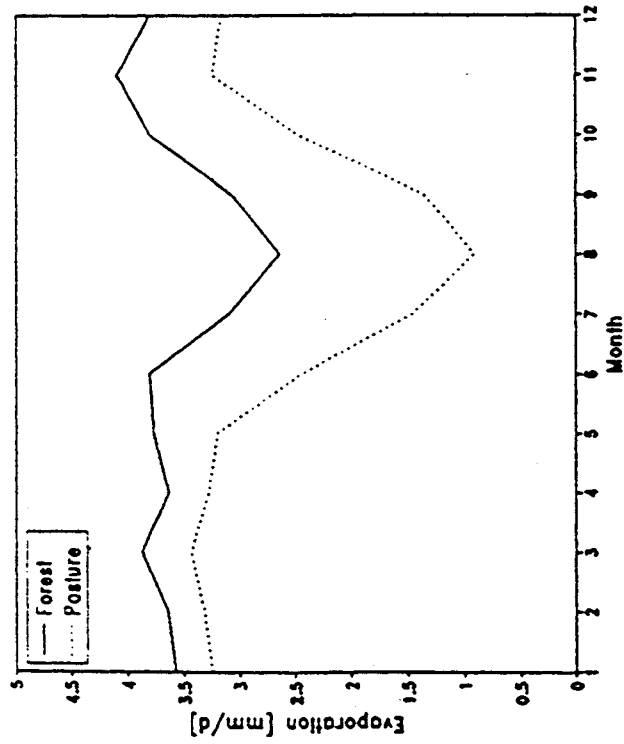
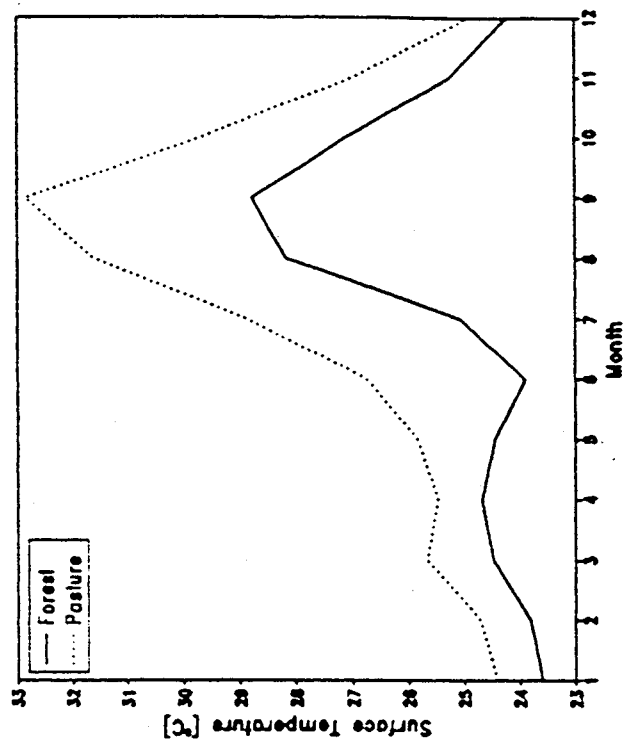
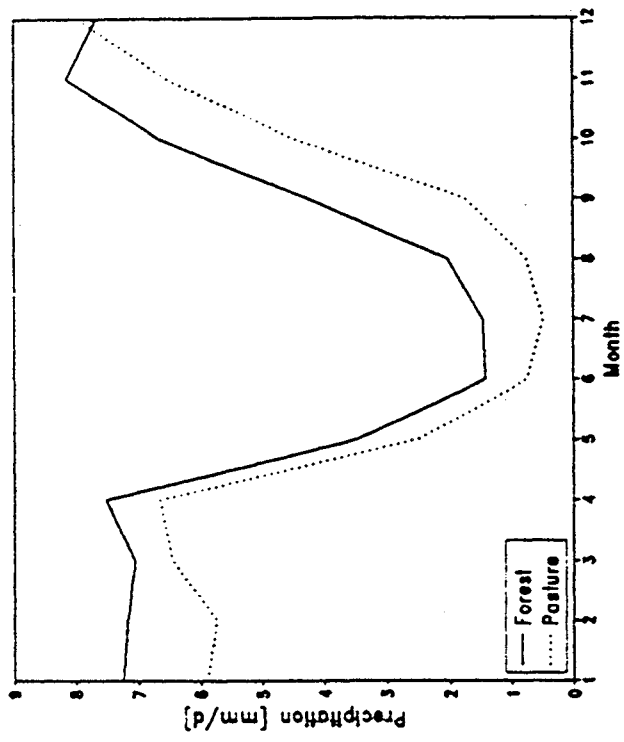
Amazon MPI model

Table 1: Initialization of the deforestation experiment.

Land surface parameters	Control	Deforestation
Vegetational roughness length [m]	0.102 - 2.000	0.026
Background albedo	0.12 - 0.23	0.18 - 0.23
Forest index	0.05 - 0.95	0.00
Vegetation ratio	0.13 - 0.96	0.11 - 0.82
Leaf area index	1.2 - 9.3	1.2 - 1.2
Max. soil moisture [m]	0.12 - 0.73	0.08 - 0.49

replace
tropical forest
by pasture

Amazon MPI model



Vegetation in the Mediterranean

Issues

Vegetation contributes to the so-called greenhouse effect

impacts of greenhouse effect

→ warming, less precipit.

Hypothesis: Was the climate during the Roman Classical time more humid because there was still a lot of vegetation

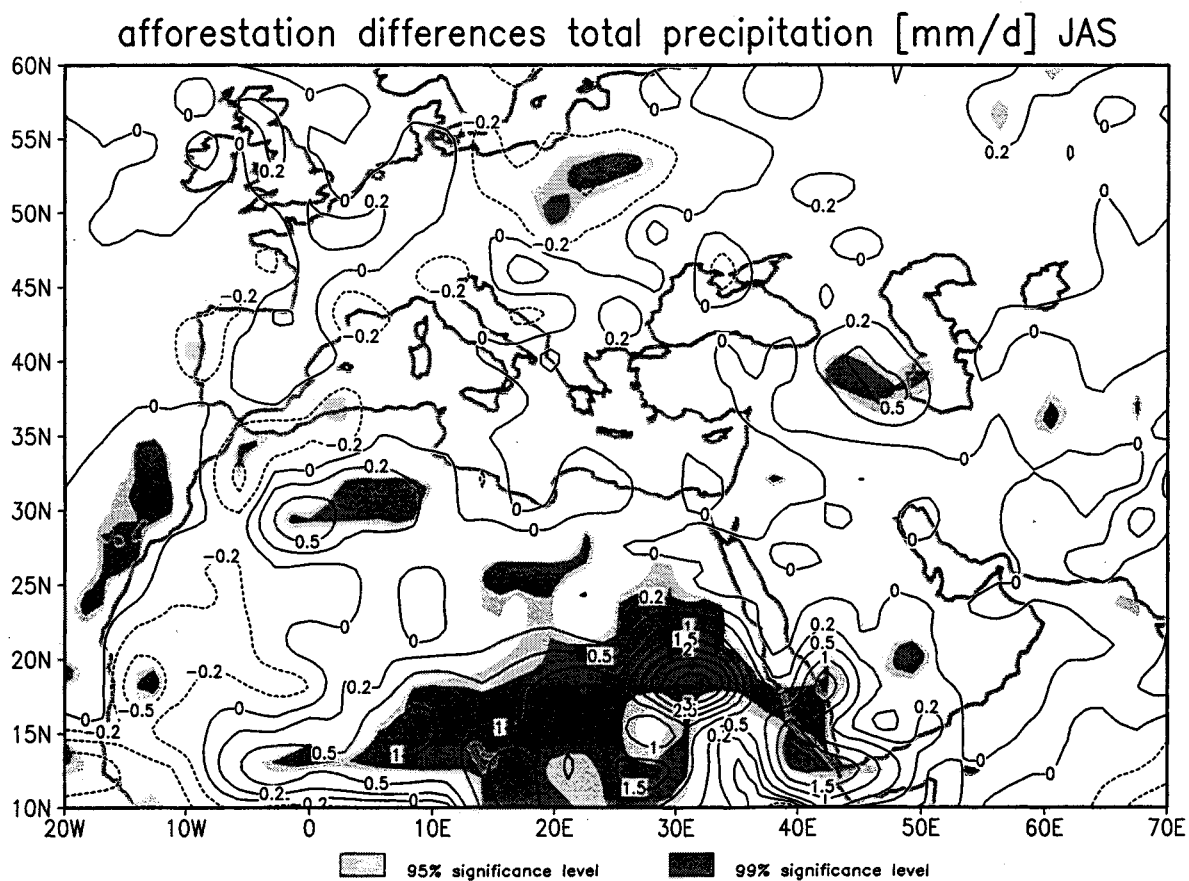
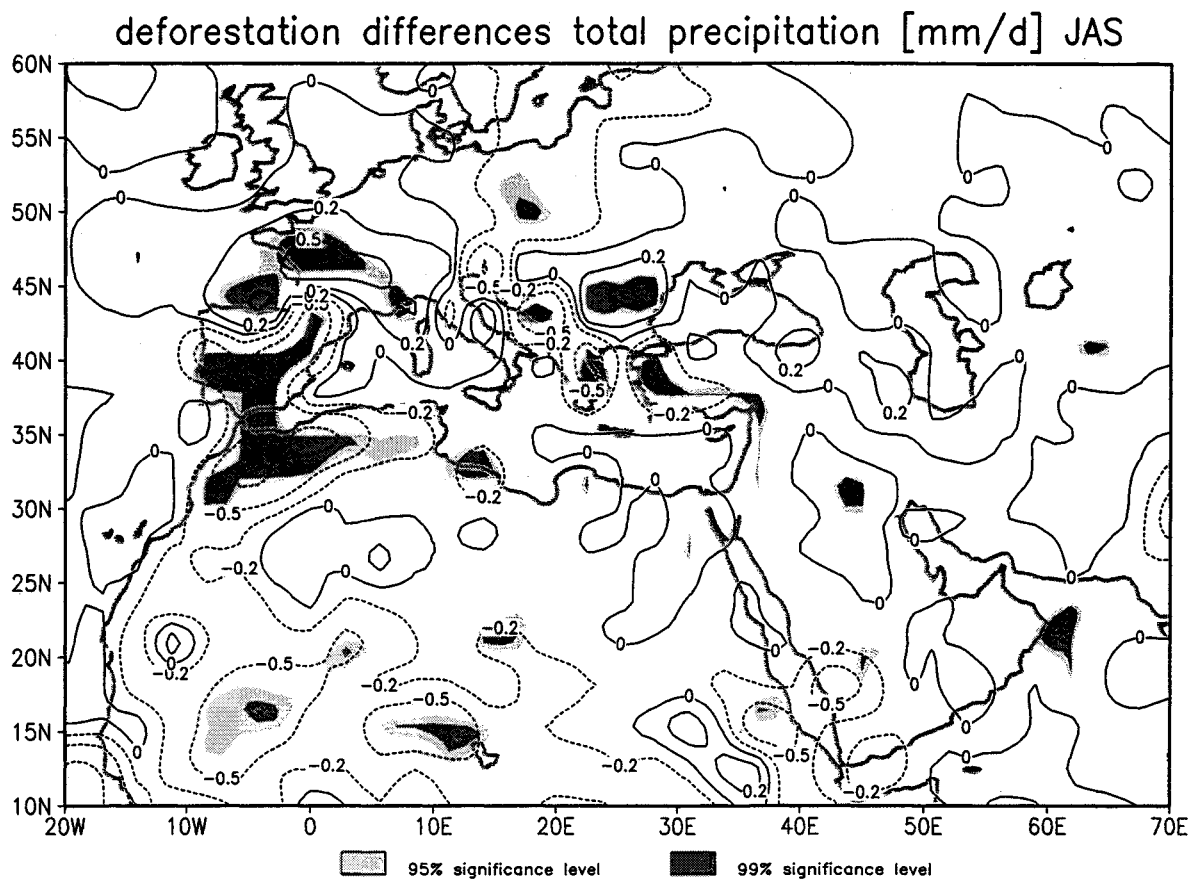
Peale and Shukla

Glob. and Planet. Ch.

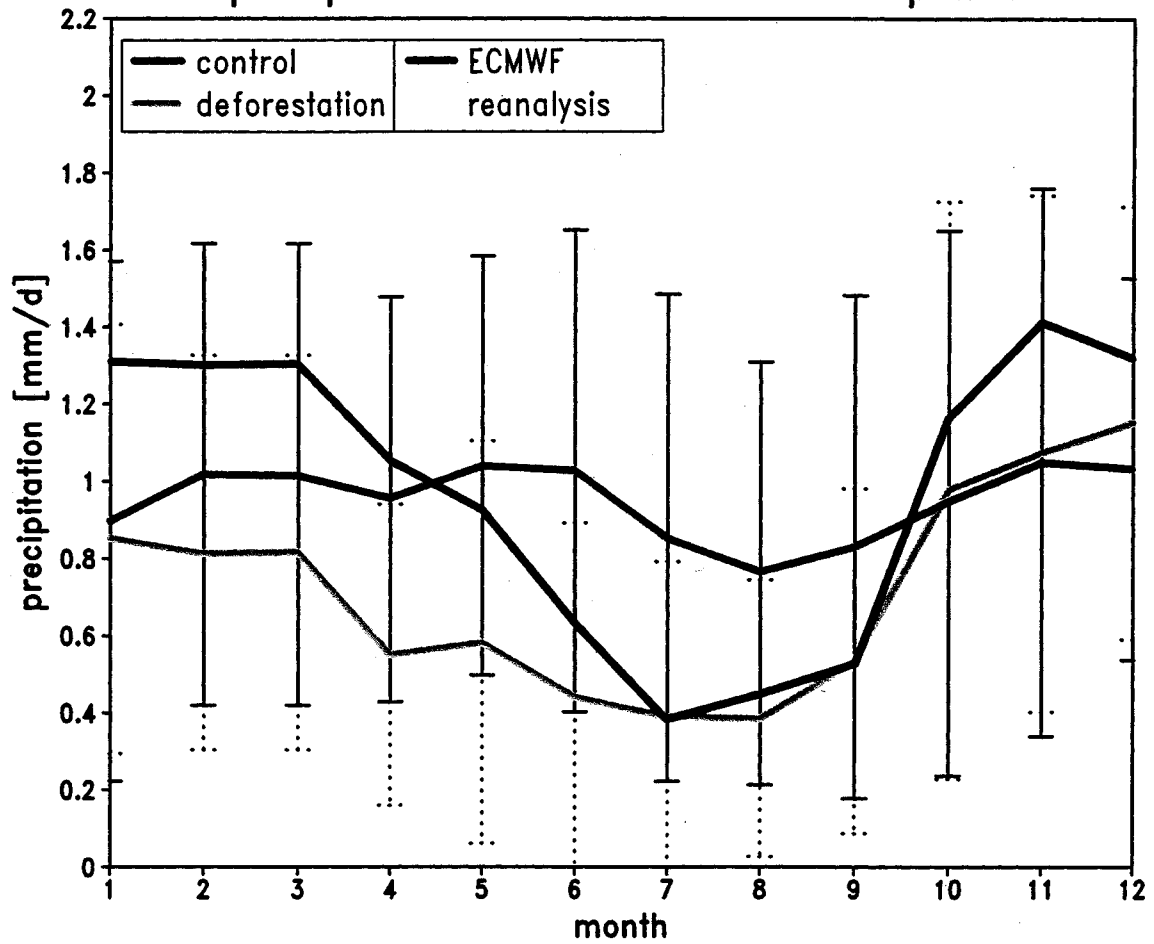
25, 2000, 185-214

Hypothesis: Water can be retained by afforestation to sustain regrowth of vegetation.

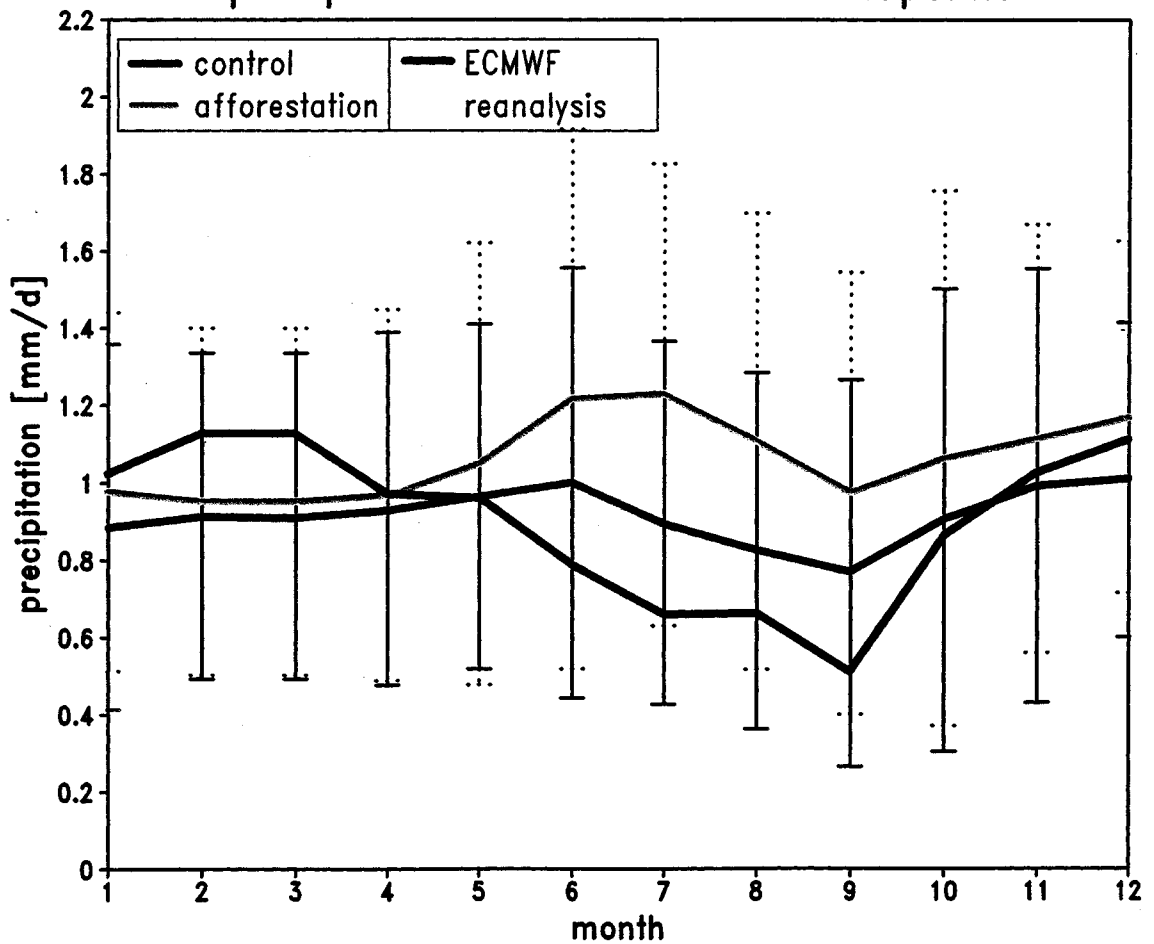
Kyoto?



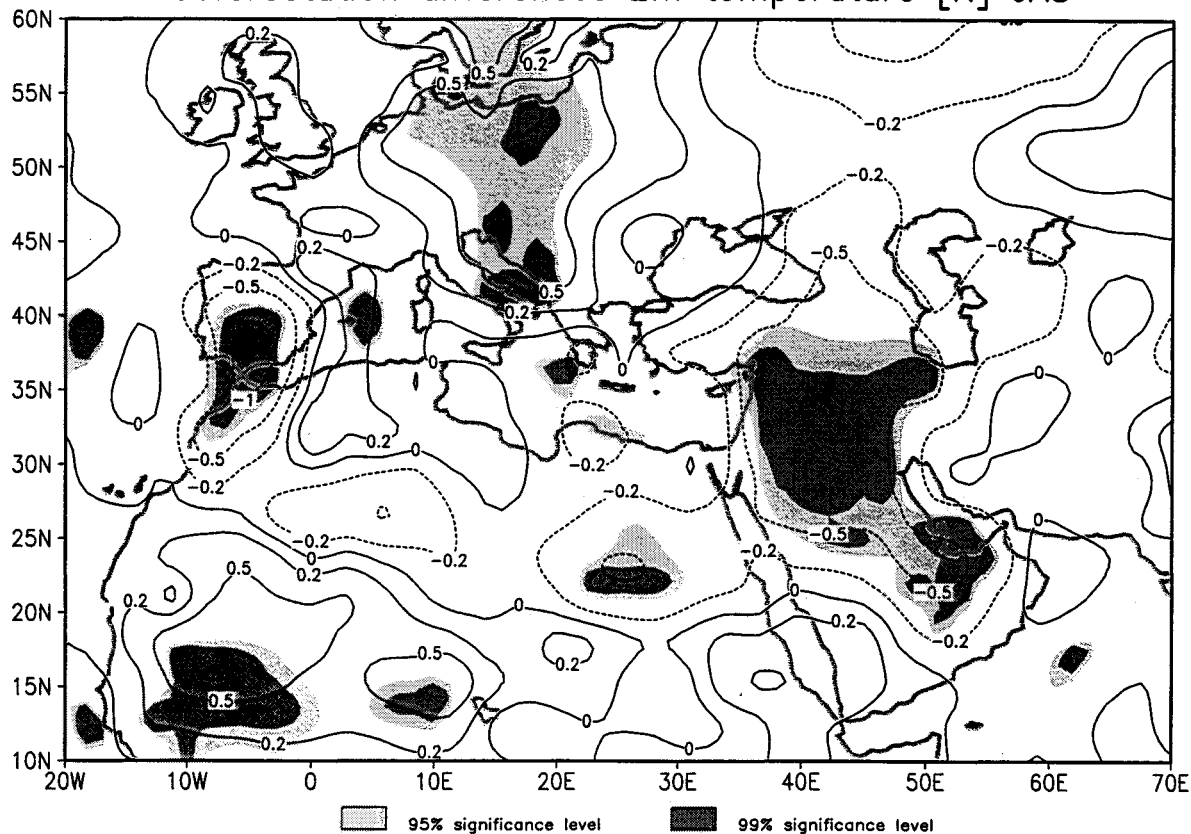
precipitation at deforestation landpoints



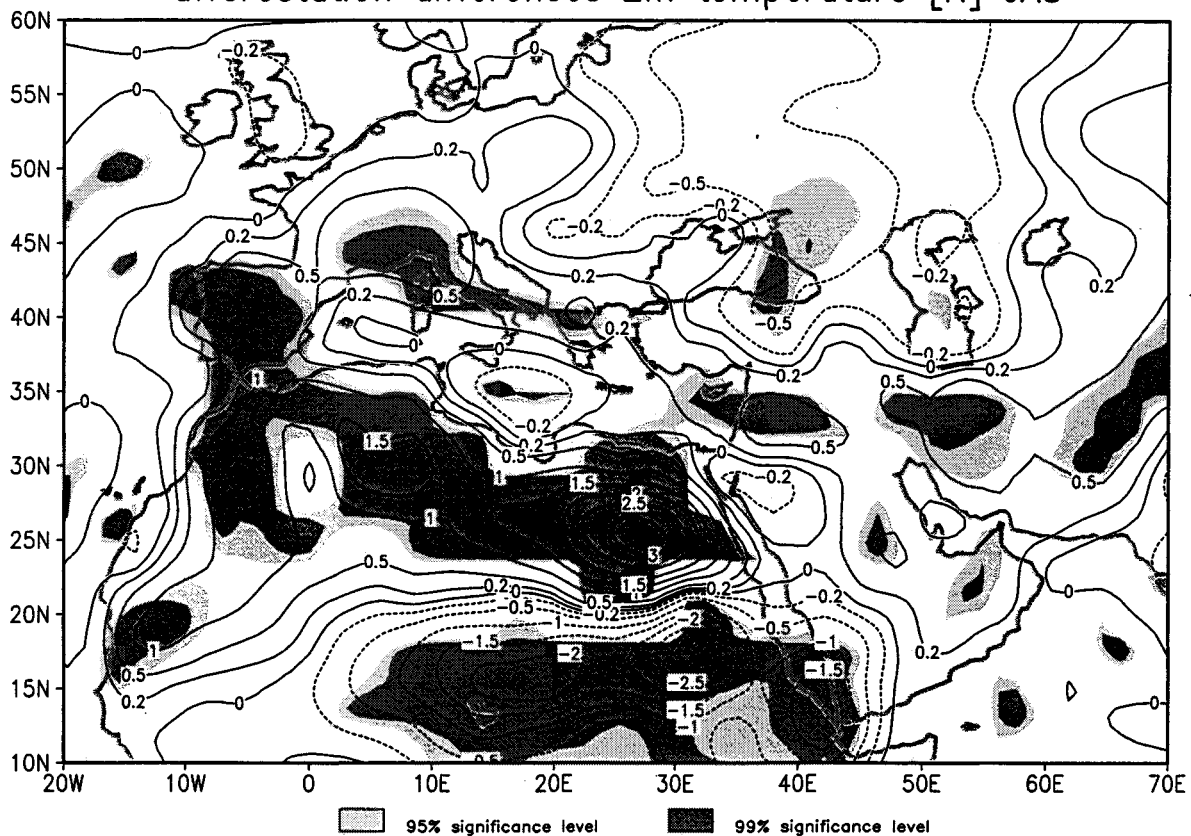
precipitation at afforestation landpoints



deforestation differences 2m temperature [K] JAS



afforestation differences 2m temperature [K] JAS



Conclusions

global model
but
no teleconnections

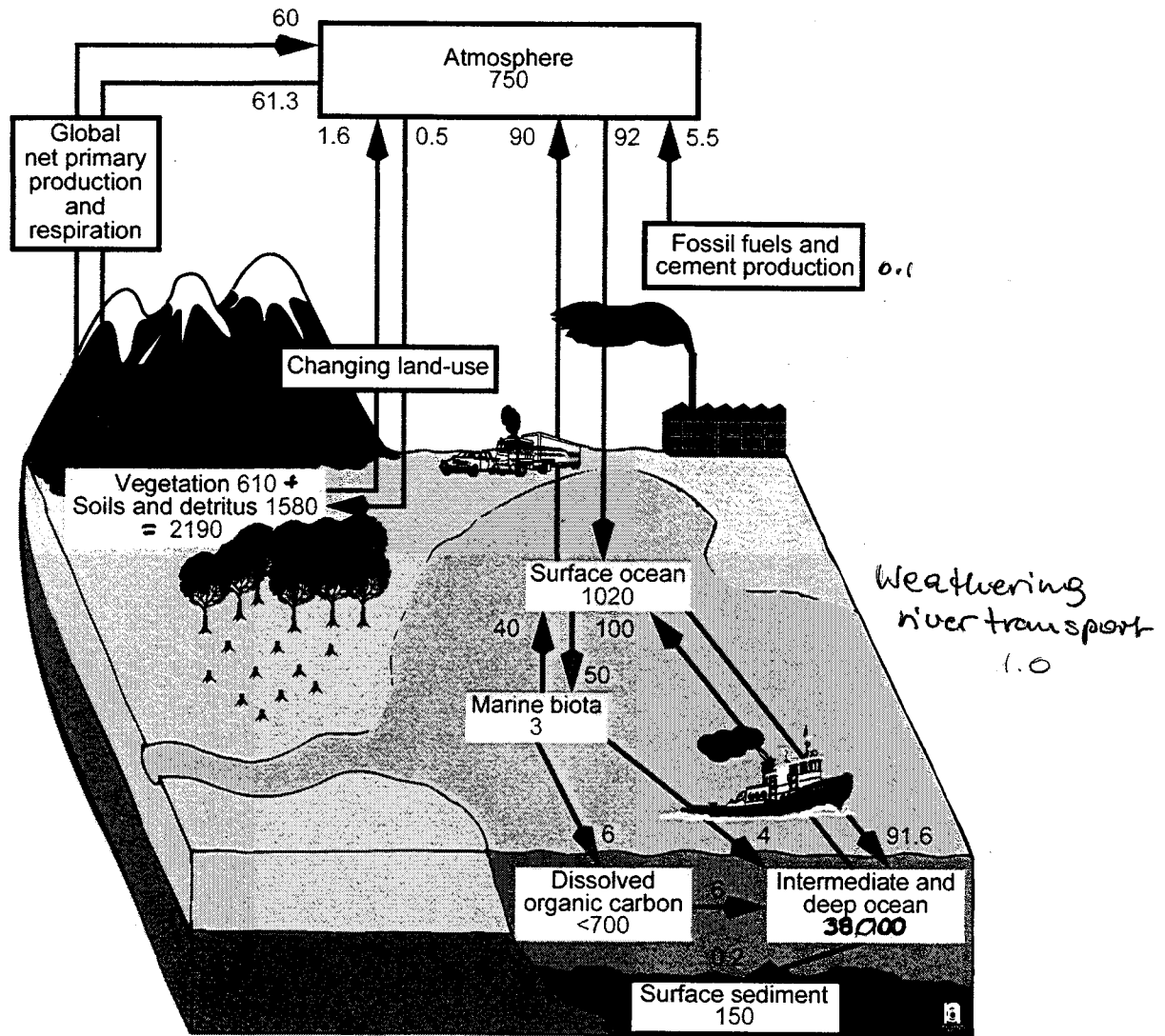
- 'study' dependence of result
- 'model' dependence of result
- 'location' dependence of result

systematic model error:
different parameteri-
zation packages

consistent physics even in global
coarse resolution model

improved local details and
more realistic variability
in time and space
→ regional models

The Global Carbon Cycle



Estimates of the various components of the global carbon cycle. Units are Gt/y (1 GT = covers the state of Hamburg with 0.5m coal) In this context ocean - atmosphere interactions and ocean processes are of crucial importance.

Volcanism < 0.1

land use change

historical cumulative losses

180-200 PgC

temperature

derived from potential vegetation maps and current vegetation estimates from satellite remote sensing.

	1980s	1990s
today :		
fossil fuel	5.4	6.3
cement prod.		

land use 1.7 ± 0.8 ?

i.e. 10-30% of total emissions of CO₂

land use flux positive: due to deforestation

negative: due to regrowth and abandoned agricult. land.

f (climate, CO₂ and N fertilization)

- Plants exchange atmospheric water vapour and carbon dioxide through leaf stomata (pores)

amount : $\frac{1}{3}$ of all atmospheric CO_2
 270 PgC / yr

- CO_2 that has exchanged oxygen with plant water oxygen carries the ^{18}O isotope signature
- During photosynthesis in the plant, CO_2 is converted to carbohydrate (gross primary production)
- $\frac{1}{2}$ of GPP is built into leaves, roots and wood.
 $\frac{1}{2}$ of GPP is given back to the atmosphere in autotrophic respiration (by plant tissue)
- Annual plant growth is net primary production (NPP) ($P_n - R_R$)
 $\sim 60 \text{ PgC / yr}$
- all carbon fixed in NPP will be returned to the atmospheric CO_2 pool by
 - heterotrophic respiration
 - natural or anthropogenic fires

- Dead biomass enters the detritus and soil organic matter pools
 - turnover time < 10 yr
 - soil organic carbon 10-100 yr
 - most soil org. carb. $>>$

Modelling

IPCC scenarios: transient simulations with a prescribed 1% increase of atmospheric CO_2 per year
(cf. SST fixed)

there is no feedback with the carbon cycle
only radiative effects and subsequent effects on water and energy cycles in ocean and atmosphere
vegetation boundary condition remains as in present-day climate

Extended coupled modelling

Cox et al. 1999

Dupresne et al. 1999

ocean carbon model (solubility biological)

terrestrial dynamic

global vegetation model
5 types (TRIFFID)

soil carbon

Carbon fluxes every 30 minutes

Update of vegetation and soil

Carbon every 10 days

dynamical evolution of land

Cover based on

Competition between types

Carbon loss from litter fall

soil carbon pool : broken up by

microorganisms

→ CO_2 returned to atmosphere

Result: land sink turns into
source as the Amazon
region vegetation dies.

(Cox et al.)

amplification of the
greenhouse effect.