## School on the Physics of Equatorial Atmosphere

 (24 September - 5 October 2001)
## Appendix to <br> Gravity Waves

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## Gravity Waves



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## Basic Wave Parametens

- The following information is required to assess the importance of gravity waves in determining the state of the tropical atmosphere
- Wave amplitudes or energy as a function of time and space
- Propagation directions in both the horizontal and vertical
- Relative importance of different wave sources
- Energy and momentum fluxes as a function of frequency and wavenumber
- How do we obtain this information from basic observations of wind, temperature and other observed quantities?


## Wave Energy

- Filter wind or temperature field to obtain wave perturbations or fluctuations
- Compute mean square amplitudes and hence energy

$$
\begin{aligned}
& U=(u, v, w) \text { where } u=\bar{u}+u^{\prime} \text { etc } \\
& E=E_{k}+E_{p} \\
& E=\frac{\mathbf{1}}{\mathbf{2}}\left(\overline{u^{\prime 2}}+\overline{v^{\prime 2}}+\overline{w^{\prime 2}}+\frac{g^{2}}{N^{2}} \overline{\hat{T}^{2}}\right) \\
& \text { where } \hat{T}=\frac{T^{\prime}}{\bar{T}}
\end{aligned}
$$

## Example from

 GPS/MET observations (Tsuda et al, 2000)High pass spatial filter with 10 km cut-off

$T$
$T^{\prime}$

$\overline{T^{\prime 2}}$
$N^{2}$

- Synoptic radiosonde observations are a valuable resource
- Fit polynomial to remove background


Jonuary 19, 1994






## $E_{\mathrm{p}}$ Climatologies




GPS/MET Nov-Feb 20-30 km
SPARC radiosonde climatollogy

## $17-24 \mathrm{~km}$

Note large equatorial amplitudes. Source effect or observational bias?

## Tropical GW



Six year time series of $E$ in 18-25 km range at Cocos Is $\left(12^{\circ} S\right)$. Note annual and $\mathrm{QBO}-$ like cycles. Source or wind filtering effects? (Vincent and Alexander, 2000)

GPS/MET observations
(Hocke and Tsuda, 2001


Note apparent geographic relation between location of enhanced fluctuations in $N, T$, and humidity (convection?).

## Definitions

To determine direction of propagation in both the vertical and horizontal use the polarization relations that relate the perturbation quantities, $u^{\prime}, v^{\prime}, w^{\prime}$, and $T^{\prime}$. Assume that we are dealing with harmonic waves in a rotating fluid

$$
\begin{aligned}
& u^{\prime}, v^{\prime}, w^{\prime}, T^{\prime} \propto \rho^{-1 / 2} e^{i(k x+l y+m z-\hat{\omega} t)} \\
& v^{\prime}=\frac{\hat{\omega} l-i f k}{\hat{\omega} k+i f l} u^{\prime} \\
& \hat{\omega}=\omega-\vec{k}_{h} \cdot \vec{U} \\
& m^{2}=\frac{N^{2}-\hat{\omega}^{2}}{\hat{\omega}^{2}-f^{2}} k_{h}^{2}-\frac{1}{4 H^{2}}
\end{aligned}
$$

$$
\hat{\omega}=\text { intrinsic frequency }
$$

$$
\omega=\text { ground - based (observed) frequency }
$$

$$
f=\text { Coriolis parameter }
$$

$$
\hat{c}=(\hat{\omega} / k, \hat{\omega} / l)=\text { intrinsic phase speed }
$$

For many wavesit is possible to assume

$$
\begin{aligned}
& m^{2} \gg \frac{1}{4 H^{2}} \\
& m= \pm \sqrt{\frac{N^{2}-\hat{\omega}^{2}}{\hat{\omega}^{2}-f^{2}}} k_{h}
\end{aligned}
$$

## Propagation Directions: Anisotropy

- Hodograph analysis

- For upward energy propagation
- Clockwise rotation in NH $(f>0)$
- Anticlockwise in SH
- Circular polarization as $\hat{\omega} \rightarrow f$
- Linear polarization as

$$
\hat{\omega} \rightarrow N
$$

## Vertical Propagation

- Rotary decomposition provides a statistical approach to hodograph analysis
$\mathbf{V}(z)=u^{\prime}(z)+i v^{\prime}(z)=a e^{-i m z}$
$\mathbf{V}(z) \leftrightarrow \Psi(m)$
Clockwise component $C(m)=\frac{1}{2} \Psi(-m) \Psi *(-m)$

$$
\mathbf{V}(z) \leftrightarrow \Psi(m)
$$

Anticlockwise component $A(m)=\frac{1}{2} \Psi(m) \Psi *(m)$
Total Energy, $E(m)=C(m)+A(m)$
In NH (SH) fraction of upward energy $\quad r_{N}=\frac{\overline{C(m)}}{\overline{E(m)}} \quad\left(r_{S}=\frac{\overline{A(m)}}{\overline{E(m)}}\right)$

## Horizontal Direction

- If axes are aligned parallel and perpendicular to major and minor axes of hodograph ellipse then

$$
\begin{aligned}
& v_{\perp}^{\prime}=-i \frac{f}{\omega} u_{\|}^{\prime} \\
& \text { Axial ratio, } A R=\frac{u_{\|}^{\prime}}{v_{\perp}^{\prime}}=\frac{\hat{\omega}}{f}
\end{aligned}
$$

- Hodograph gives
- direction of propagation (with $180^{\circ}$ ambiguity)
- intrinsic frequency.

Statistical approach is better.

## Stokes Parameters

- Note similarity between gravity wave and electromagnetic wave
- Both are transverse oscillations that are in general elliptically polarised
- Stokes parameters are relations that determine sense and degree of polarisation

$$
\begin{array}{ll}
I=\overline{u^{\prime 2}}+\overline{v^{\prime 2}} & I=\text { total intensity } \\
D=\overline{u^{\prime 2}}-\overline{v^{\prime 2}} & D=\text { intensity difference } \\
P=\overline{\mathbf{2} u^{\prime} v^{\prime} \cos \delta} & P=\text { Linear polarization parameter } \\
Q=\overline{\mathbf{2} u^{\prime} v^{\prime} \sin \delta} & \mathbf{Q}=\text { circular polarization parameter }
\end{array}
$$

## Stokes Parameters II

- $Q$ gives sense of rotation
$-Q>0 \rightarrow$ Anticlockwise polarisation
$-Q<0 \rightarrow$ Clockwise polarisation
- Degree of polarisation
$d=\frac{\left(D^{2}+P^{2}+Q^{2}\right)^{1 / 2}}{I}$
$d=\mathbf{1} \Rightarrow$ perfect polarisation
$d=0 \Rightarrow$ unpolarised


## Polarisation Ellipse

- Polarisation ellipse parameters

$$
\begin{aligned}
& \delta=\boldsymbol{\operatorname { t a n }}^{-1}\left(\frac{Q}{P}\right) \\
& \mathbf{2 \phi}=\boldsymbol{\operatorname { t a n }}^{-1}\left(\frac{P}{D}\right) \\
& A R=\boldsymbol{\operatorname { c o t }} \varepsilon \text { where } 2 \varepsilon=\boldsymbol{\operatorname { s i n }}^{-1}\left(\frac{Q}{d I}\right)
\end{aligned}
$$

Apply in frequency domain using radar winds data acquired at a single height

Apply in vertical wavenumber domain using rocket or radiosonde wind profiles

## Horizontal Directions

- Require extra information to resolve $180^{\circ}$ directional ambiguity from hodograph
- Use either vertical velocity or temperature
$w^{\prime}=-\frac{k_{h}}{m} u_{\|}^{\prime}$

$w^{\prime}$ in - phase with $u^{\prime}$ for $k_{h}>0$ and $m<\mathbf{0} \Rightarrow \overline{u_{\|}^{\prime} w^{\prime}}>\mathbf{0}$
$w^{\prime}$ out - of - phase with $u^{\prime}$ for $k_{h}<0$ and $m<0 \Rightarrow \overline{u_{\|}^{\prime} w^{\prime}}<0$


## Horizontal Directions II

- Temperature polarisation relation

$$
\hat{T}^{\prime}=\mp i \frac{N^{2}}{g \hat{\omega}} \frac{k_{h}}{m} u_{\|}^{\prime}
$$

- sign chosen when $k m>0$
+ sign chosen when $k m<0$


Note $90^{\circ}$ phase shift between $T$ and $u$
$\overline{u_{\|}^{\prime} w^{\prime}}=\frac{\hat{\omega} g}{N^{2}} \overline{u_{\|}^{\prime} \hat{T}_{+90}^{\prime}}$ where $\hat{T}_{+90}^{\prime}=i \hat{T}^{\prime}$
Sign of $u_{\|}^{\prime} \hat{T}_{+90}^{\prime}$ determines direction of propagation


## Momentum Fluxes

 $u^{\prime} w^{\prime}, v^{\prime} w^{\prime}$- Important quantities, but difficult to measure directly ( $w^{\prime} \ll u^{\prime}, v^{\prime}$ )
- Indirect estimates using temperature require knowledge of $\hat{\boldsymbol{\omega}}$

If $\bar{U}=\boldsymbol{0}$ then $\hat{\omega}=\boldsymbol{\omega}$, the ground -based frequency

## Radar Measurements

$u^{\prime} w^{\prime}, \nu^{\prime} w^{\prime}$

- Dual-beam Doppler technique
- Measure radial velocity at range, $R$, along beams 1 and 2

$$
\begin{aligned}
& v_{1}(\theta, R)=u^{\prime} \sin \theta+w^{\prime} \cos \theta \\
& v_{2}(-\theta, R)=u^{\prime} \sin \theta-w^{\prime} \cos \theta \\
& v_{1}^{2}(\theta, R)=u^{\prime 2} \sin ^{2} \theta+2 u^{\prime} w^{\prime} \cos \theta \sin \theta+w^{\prime 2} \cos ^{2} \theta \\
& v_{2}^{2}(-\theta, R)=u^{\prime 2} \sin ^{2} \theta-2 u^{\prime} w^{\prime} \cos \theta \sin \theta+w^{\prime 2} \cos ^{2} \theta
\end{aligned}
$$



## Subtract and rearrange

$$
\overline{u^{\prime} w^{\prime}}(z)=\frac{\overline{v_{1}^{2}(\theta, R)}-\overline{v_{2}^{2}(-\theta, R)}}{2 \sin 2 \theta}
$$

## Strengths:

- Robust
- Direct

Limitations:

- Require accurate beam angles
- Small differences


## Spectra

- Frequency and wavenumber spectra are a convenient way of summarizing wave information
- Suggestion that there is a "universal" wave spectrum
- Spectral shapes and amplitudes set by instability processes
- Many different theories for producing spectral features
- Observations show that
- Frequency spectra derived from ground-based radar and lidar observations of horizontal winds show $S_{\mathrm{u}}(\omega) \propto \omega^{-\mathrm{p}}$ with $p \sim 5 / 3$
- Recent constant pressure balloon observations provide first intrinsic frequency spectra

$$
S_{u}(\hat{\omega}) \propto \hat{\omega}^{-2}
$$

- Vertical wavenumber spectra derived from radiosonde, rocket, and radar observations of winds and temperatures show $S_{\mathbf{u}, \mathrm{T}}(\mathbf{m}) \propto m^{-\mathrm{t}}$ with $t \sim \mathbf{3}$ for large $m$


## Spectra: Examples



Modified Desaubies spectrum

$$
\begin{aligned}
& F(\mu)=\frac{\mu^{s}}{\mathbf{1}+\mu^{s+t}} \\
& \mu \equiv \frac{m}{m_{*}}, t \sim \mathbf{3}, s \sim \mathbf{1}(\mathbf{?})
\end{aligned}
$$

## Intrinsic Frequency Spectra

- Trans-Pacific constantpressure balloon flights
- 19-20 km float altitude
- Zonal, meridional, vertical velocity spectra





Circumpolar spectra

## Observational Selection

- All observational techniques have some bias in which part of the wave spectrum they can observe
- No one technique can cover the complete wave spectrum
- Important to recognize and quantify the possible biases



## Observational filters for MLS, rockets, (GPS/MET), radiosondes

Modelling of observations helps with interpretation
e.g. Alexander and Vincent (2000); McLandress et al (2000)

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Radiosonde observations constrain model



Gravity-wave driven force

Force estimate from "best-fit" model

Force estimated without allowing for observational selection

(b)


Understanding MLS



