

the **abdus salam** international centre for theoretical physics

SMR/1328/3 bis

School on the Physics of Equatorial Atmosphere

(24 September - 5 October 2001)

Appendix to

Gravity Waves

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Gravity Waves International School on the Physics of the Equatorial Atmosphere and Ionosphere **ICTP** Trieste September 2001 ÎESE **R. A. Vincent**

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International School on Equatorial Atmosphere: Gravity Waves

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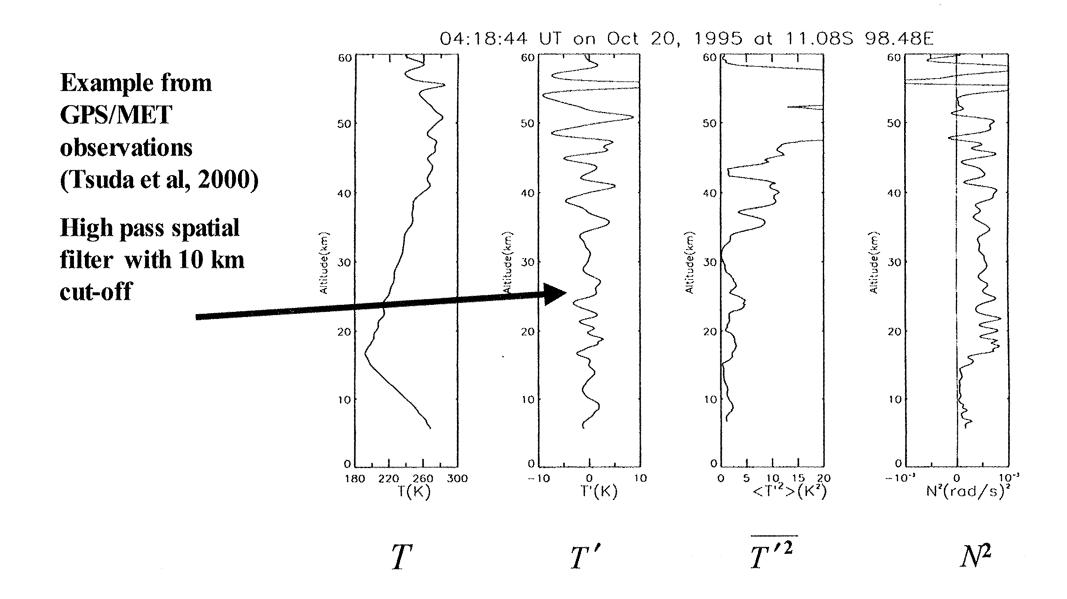
Basic Wave Parameters

- The following information is required to assess the importance of gravity waves in determining the state of the tropical atmosphere
 - Wave amplitudes or energy as a function of time and space
 - Propagation directions in both the horizontal and vertical
 - Relative importance of different wave sources
 - Energy and momentum fluxes as a function of frequency and wavenumber
- How do we obtain this information from basic observations of wind, temperature and other observed quantities?

Wave Energy

- Filter wind or temperature field to obtain wave perturbations or fluctuations
- Compute mean square amplitudes and hence energy

 $U = (u, v, w) \text{ where } u = \overline{u} + u' \text{ etc}$ $E = E_k + E_p$ $E = \frac{1}{2} \left(\overline{u'^2} + \overline{v'^2} + \overline{w'^2} + \frac{g^2}{N^2} \overline{\hat{T}^2} \right)$ where $\hat{T} = \frac{T'}{\overline{T}}$

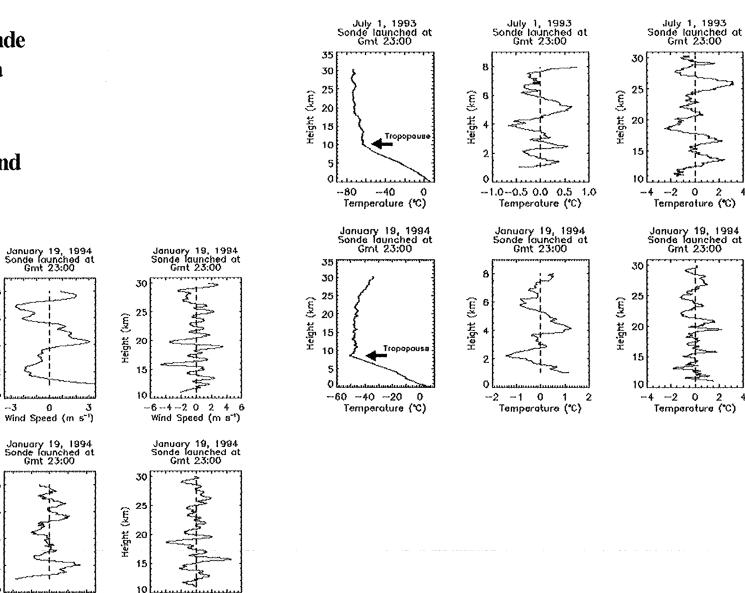


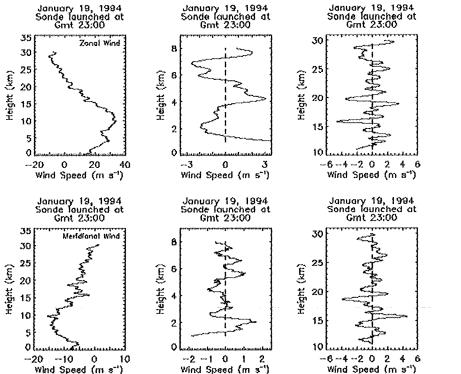
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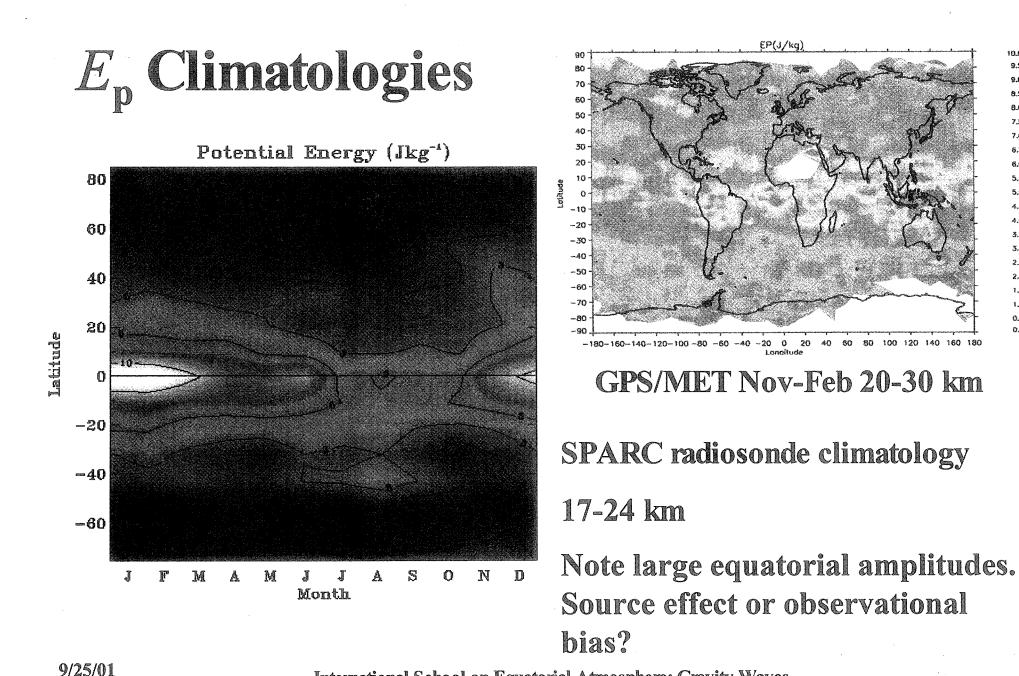
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- Synoptic radiosonde ۲ observations are a valuable resource
- Fit polynomial to . remove background







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6

(J/kg

9.5 9.0 8.5 8.0 7.5 7.0

6.5 6.0

5.5

5.0 4.5 4.0

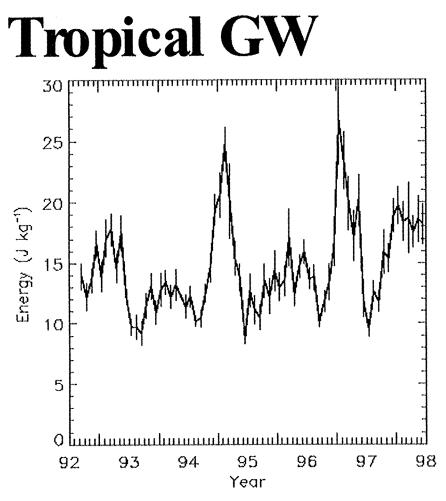
3.5 3.0 2.5

2,0

1.5

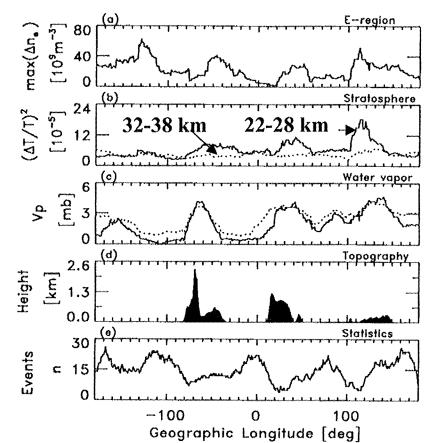
1.0

0.5



Six year time series of E in 18-25 km range at Cocos Is (12°S). Note annual and QBOlike cycles. Source or wind filtering effects? (Vincent and Alexander, 2000)

GPS/MET observations (Hocke and Tsuda, 2001



Note apparent geographic relation between location of enhanced fluctuations in *N*, *T*, and humidity (convection?).

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Definitions

To determine direction of propagation in both the vertical and horizontal use the polarization relations that relate the perturbation quantities, u', v', w', and T'. Assume that we are dealing with harmonic waves in a rotating fluid

$$u', v', w', T' \propto \rho^{-1/2} e^{i(kx+ly+mz-\delta t)}$$

$$v' = \frac{\hat{\omega}l - ifk}{\hat{\omega}k + ifl} u'$$

$$\hat{\omega} = \omega - \vec{k}_h \cdot \vec{U}$$

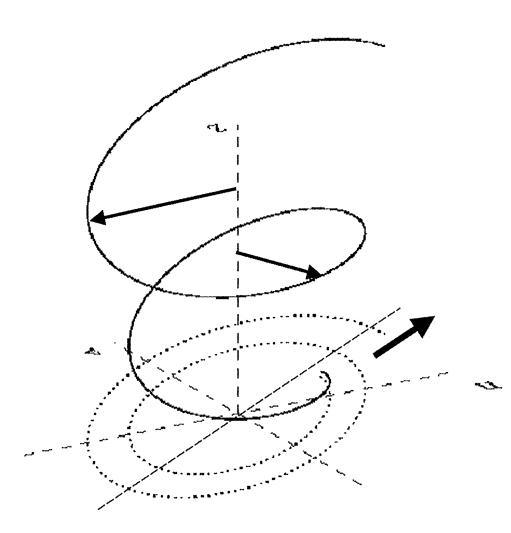
$$m^2 = \frac{N^2 - \hat{\omega}^2}{\hat{\omega}^2 - f^2} k_h^2 - \frac{1}{4H^2}$$

 $\hat{\omega}$ = intrinsic frequency ω = ground - based (observed) frequency f = Coriolis parameter $\hat{c} = (\hat{\omega}/k, \hat{\omega}/l)$ = intrinsic phase speed

For many waves it is possible to assume

$$m^{2} >> \frac{1}{4H^{2}}$$
$$m = \pm \sqrt{\frac{N^{2} - \hat{\omega}^{2}}{\hat{\omega}^{2} - f^{2}}} k_{h}$$

Propagation Directions: Anisotropy



- Hodograph analysis
- For upward energy propagation
 - Clockwise rotation in NH (f > 0)
 - Anticlockwise in SH
- Circular polarization as $\hat{\omega} \rightarrow f$
- Linear polarization as

 $\hat{\omega} \to N$

Vertical Propagation

• Rotary decomposition provides a statistical approach to hodograph analysis

 $V(z) = u'(z) + iv'(z) = ae^{-imz}$ $V(z) \leftrightarrow \Psi(m)$

Clockwise component $C(m) = \frac{1}{2} \Psi(-m) \Psi^*(-m)$

 $V(z) \leftrightarrow \Psi(m)$

Anticlockwise component $A(m) = \frac{1}{2}\Psi(m)\Psi^*(m)$ Total Energy, E(m) = C(m) + A(m)

In NH (SH) fraction of upward energy

$$r_N = \frac{\overline{C(m)}}{\overline{E(m)}} \quad \left(r_S = \frac{\overline{A(m)}}{\overline{E(m)}}\right)$$

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Horizontal Direction

• If axes are aligned parallel and perpendicular to major and minor axes of hodograph ellipse then

$$v'_{\perp} = -i\frac{f}{\hat{\omega}}u'_{\parallel}$$

Axial ratio, $AR = \frac{u'_{\parallel}}{v'_{\perp}} = \frac{\hat{\omega}}{f}$

- Hodograph gives
 - direction of propagation (with 180° ambiguity)
 - intrinsic frequency.

Statistical approach is better.

Stokes Parameters

- Note similarity between gravity wave and electromagnetic wave
- Both are transverse oscillations that are in general elliptically polarised
- Stokes parameters are relations that determine sense and degree of polarisation

$$=\overline{u'^2}+\overline{v'^2}$$

Ţ

I =total intensity

$$D = \overline{u'^2} - \overline{v'^2}$$

 $P = \overline{2u'v'\cos\delta}$

 $Q = \overline{2u'v'\sin\delta}$

- *P* = Linear polarization parameter
- $\mathbf{Q} = \mathbf{circular polarization parameter}$

Stokes Parameters II

- Q gives sense of rotation
 - $-Q > \mathbf{0} \rightarrow$ Anticlockwise polarisation
 - $-Q < \mathbf{0} \rightarrow \mathbf{Clockwise}$ polarisation
- Degree of polarisation

$$d = \frac{(D^2 + P^2 + Q^2)^{1/2}}{I}$$
$$d = 1 \Rightarrow \text{ perfect polarisation}$$

 $d = \mathbf{0} \Rightarrow$ unpolarised

Polarisation Ellipse

• Polarisation ellipse parameters

$$\delta = \tan^{-1} \left(\frac{Q}{P} \right)$$

$$2\phi = \tan^{-1} \left(\frac{P}{D} \right)$$

$$AR = \cot \varepsilon \text{ where } 2\varepsilon = \sin^{-1} \left(\frac{Q}{dI} \right)$$

Apply in frequency domain using radar winds data acquired at a single height

Apply in vertical wavenumber domain using rocket or radiosonde wind profiles

Horizontal Directions

- Require extra information to resolve 180° directional ambiguity from hodograph
- Use either vertical velocity or temperature

$$w' = -\frac{k_h}{m}u'_{\parallel}$$

w' in - phase with u' for $k_h > 0$ and $m < 0 \Rightarrow \overline{u'_{\parallel}w'} > 0$ w' out - of - phase with u' for $k_h < 0$ and $m < 0 \Rightarrow \overline{u'_{\parallel}w'} < 0$

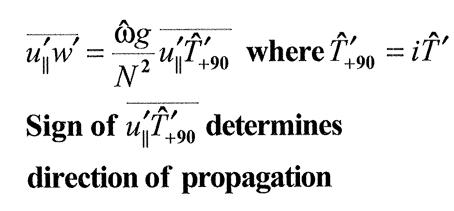
Horizontal Directions II

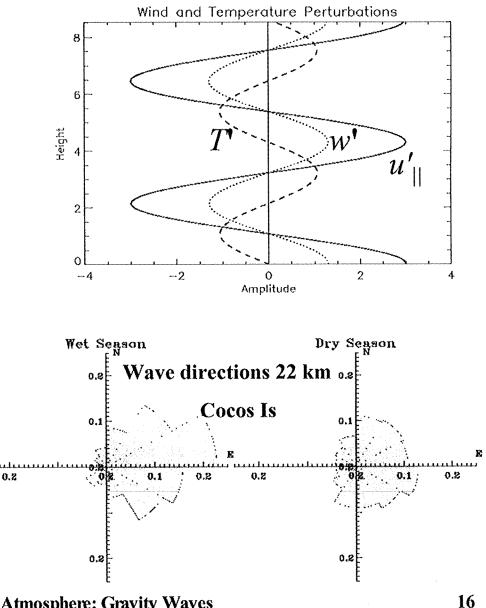
• Temperature polarisation relation

 $\hat{T}' = \mp i \frac{N^2}{g\hat{\omega}} \frac{k_h}{m} u'_{\parallel}$

 $- \operatorname{sign} \operatorname{chosen} \operatorname{when} km > \mathbf{0}$ $+ \operatorname{sign} \operatorname{chosen} \operatorname{when} km < \mathbf{0}$

Note 90° phase shift between T and u



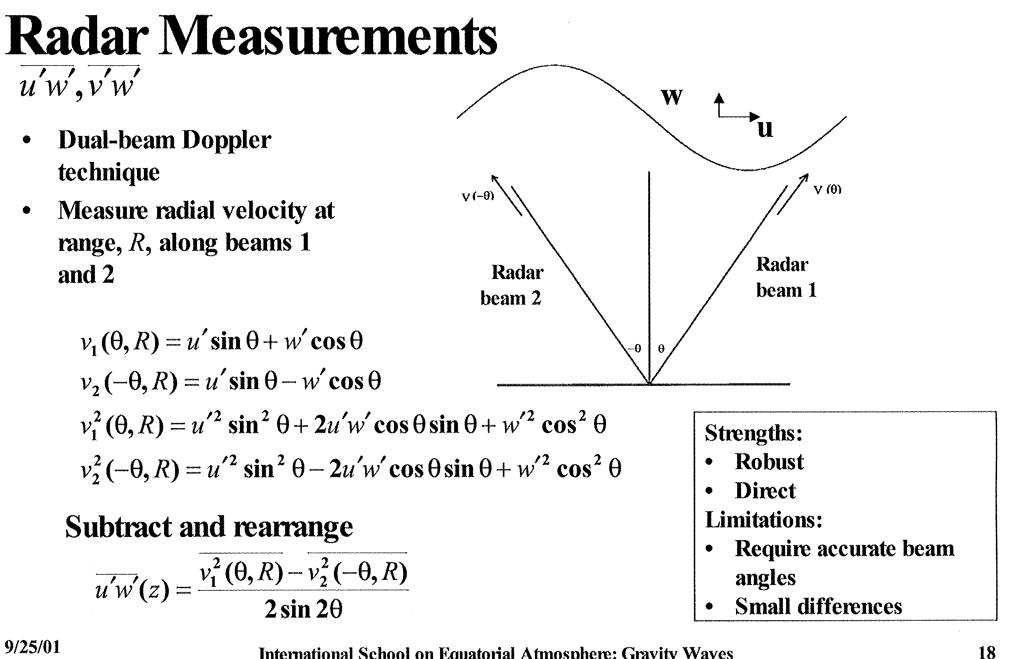


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Momentum Fluxes $\vec{u'w'}, \vec{v'w'}$

- Important quantities, but difficult to measure directly (w' << u',v')
- Indirect estimates using temperature require knowledge of $\widehat{\omega}$

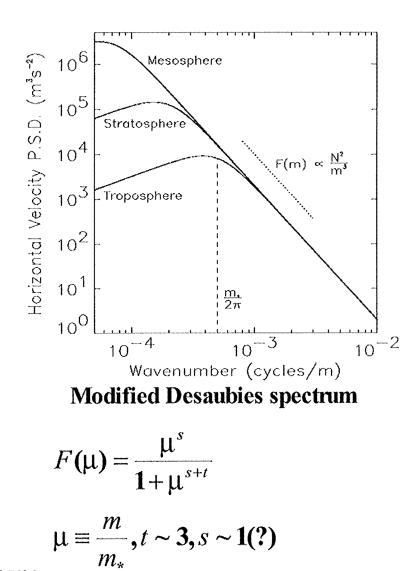
If $\overline{U} = 0$ then $\hat{\omega} = \omega$, the ground - based frequency

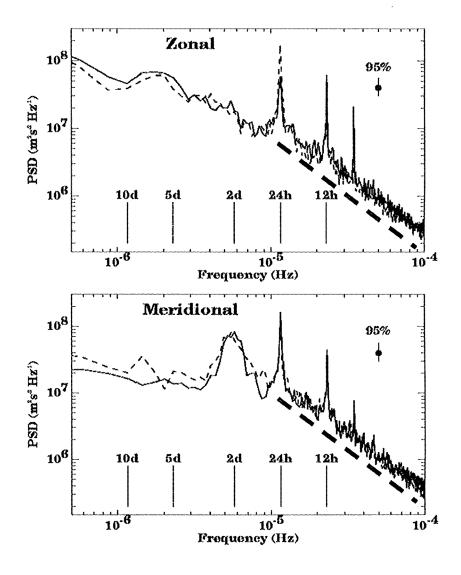


Spectra

- Frequency and wavenumber spectra are a convenient way of summarizing wave information
- Suggestion that there is a "universal" wave spectrum
- Spectral shapes and amplitudes set by instability processes
- Many different theories for producing spectral features
- Observations show that
 - Frequency spectra derived from ground-based radar and lidar observations of horizontal winds show $S_u(\omega) \propto \omega^{-p}$ with $p \sim 5/3$
 - Recent constant pressure balloon observations provide first intrinsic frequency spectra $S_u(\hat{\omega}) \propto \hat{\omega}^{-2}$
 - Vertical wavenumber spectra derived from radiosonde, rocket, and radar observations of winds and temperatures show $S_{u,T}(m) \propto m^{-t}$ with $t \sim 3$ for large m

Spectra: Examples





Equatorial mesospheric frequency spectra

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Intrinsic Frequency Spectra

Trans-Pacific constantpressure balloon flights 19-20 km float altitude Ж Zonal, meridional, vertical velocity spectra 10-3 100.00 flight 3 102 10 10 $\langle v'' \rangle = 3.8 \text{ m}^{-1} \text{ s}$ Ĵ m² s²/(cyc d") 100 s*//cyc m² s*/(cyc m² s²/(cyc d¹) 100, 10 10 10-6 10-2 0.10 10-4 0.0 10-4 0.10 1.00 10.00 100.00 10.00 100.00 0.01 0.10 1.00 0.01 cyc d" cyc c" 0.1 1.0 10.0 100.0 cyc d*

Circumpolar spectra

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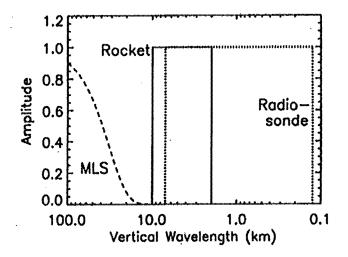
10¹

100

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Observational Selection

- All observational techniques have some bias in which part of the wave spectrum they can observe
- No one technique can cover the complete wave spectrum
- Important to recognize and quantify the possible biases



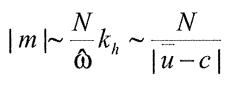
Observational filters for MLS, rockets, (GPS/MET), radiosondes

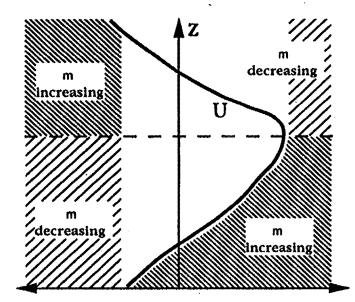
Modelling of observations helps with interpretation

e.g. Alexander and Vincent (2000); McLandress et al (2000)

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Simple dispersion relation

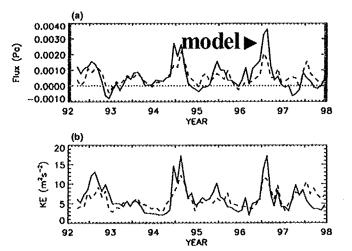




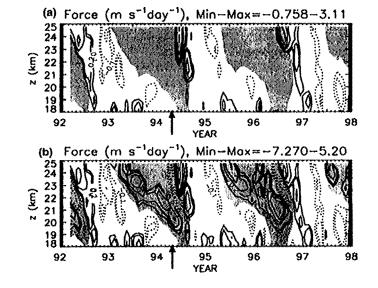
U or c (m/s) Waves move in or out of each "window" as *m* changes due to heightvarying background winds

Modelling of observations helps with interpretation

e.g. Alexander and Vincent (2000); McLandress et al (2000)



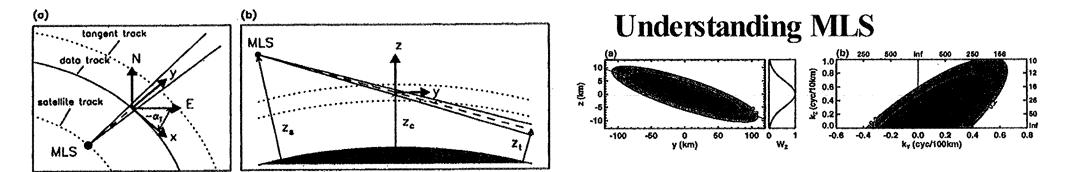
Radiosonde observations constrain model



Gravity-wave driven force

Force estimate from "best-fit" model

Force estimated without allowing for observational selection



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