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UNITED NATIONS EDUCATIONAL, SCIENTIFIC AND CULTURAL ORGANIZATION



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SMR/107 - 19

WORKSHOP ON PATTERN RECOGNITION AND ANALYSIS OF SEISMICITY

(5 - 16 December 1983)

STOCHASTIC MODELING

L. KNOPOFF

STOCHASTIC MODELING:

II-1

Physical Model of Earthquake Process
with random elements.

- 1. Branching Process
- 2. Kolmogorov/Feller Model with continuous states.
- 3. Two-dimensional Finite grid models.

To undisciplined statistics we try to add a little physics.

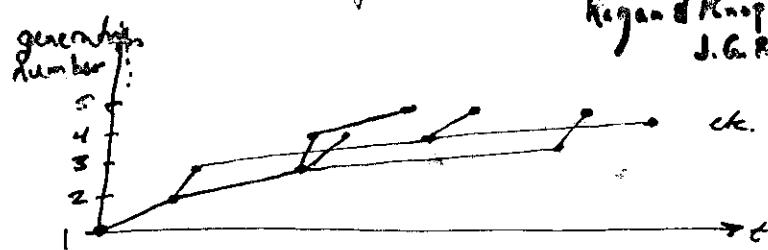
This makes statistical analysis difficult

↓
computer (Monte Carlo)

→ We need a model!

1. Critical Branching Process

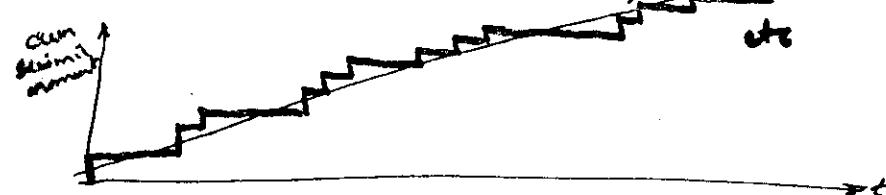
Ref. II-2
Kagan & Knopoff.
J.G.R. 86 (Aug) 1981.



Prob. that one event gives "birth" to another event in time interval dt is

$$\begin{aligned} P(t)dt &= 0 \quad \text{at } t < t_0 \\ &= (1-K)e^{kt} \frac{dt}{e^{kt}} \quad t > t_0 \end{aligned}$$

All elementary shocks assumed equal



If $K=0$ then each event gives rise to one offspring (critical mass)
on the average in infinite time

If $K < 1$ the process is subcritical.

for critical process mean free running sum total no. of shocks

may be ∞ .

For subcritical process, finite no. of shocks.

* Self-similarity!

If $\theta \sim 0.5$ (shallow aq.)
 ~ 0.8 (intermed. aq.)

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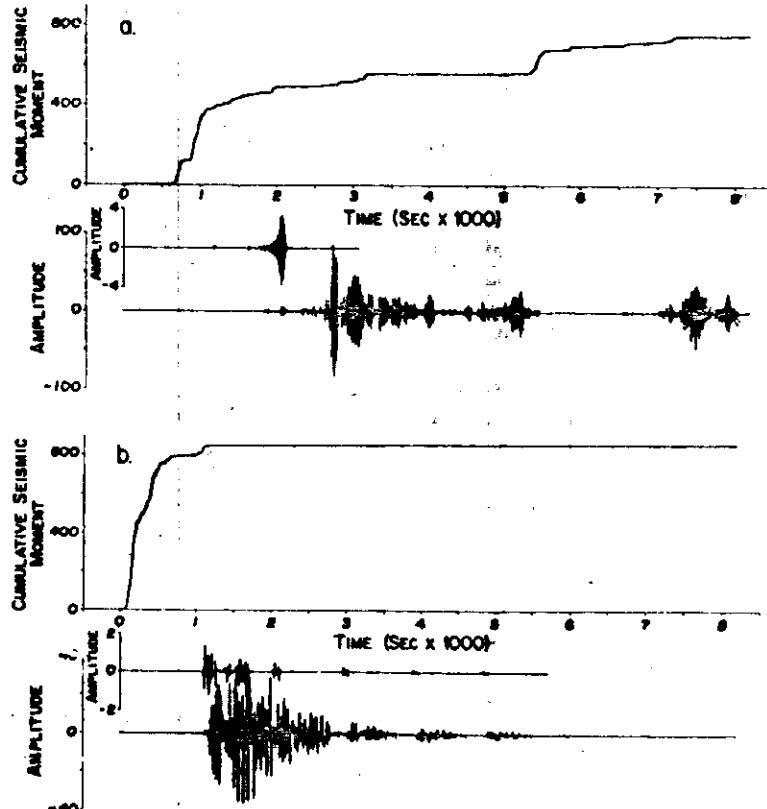
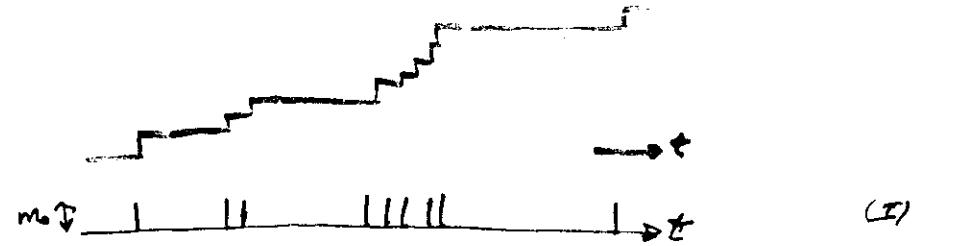


Fig. 3. Simulated source-time functions and seismograms for shallow (a) and intermediate (b) earthquake sources. The upper trace in each plot is a synthetic source-time function. The middle plot is a theoretical seismogram described in the text, and the lower trace is a convolution of the derivative of source-time function with the theoretical seismogram.

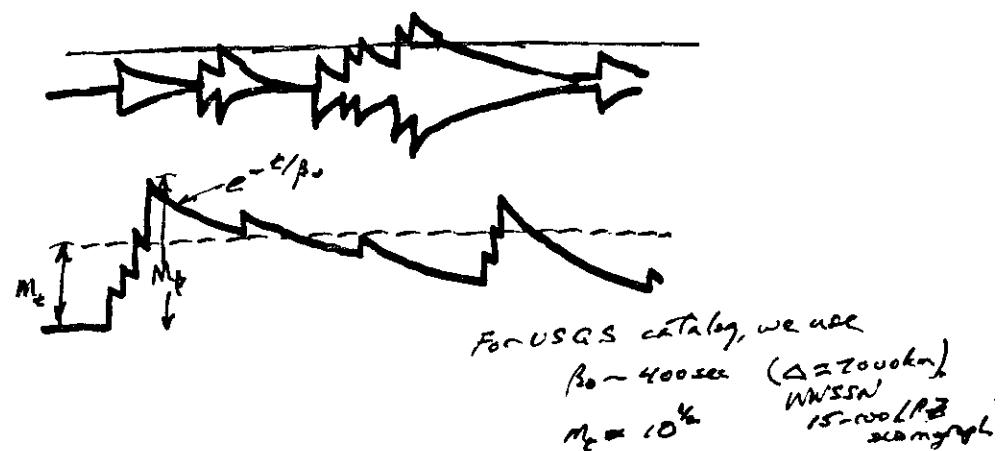


Identify M_0 with threshold ~~magnitude~~ moment of detectability of network (i.e. threshold of homogeneity).

$$\text{If } M_0 = 10^5, \quad t_0 = 2 \text{ sec (NBS catalog)}$$

$$M_0 = 10^4, \quad t_0 = 10^{-2} \text{ sec (USGS catalog)} \\ \text{California}$$

Now convolve (I) with Green's function for elastic medium
(Synthetic Seismogram)



For USGS catalog, we use
 $\beta_0 = 400 \text{ sec}$ ($\Delta = 200 \text{ km}$)
 $M_p = 10^{14}$ WISSA
 $15-100 \text{ Hz}$
 $10-100 \text{ sec}$

Dimensionless Variables:

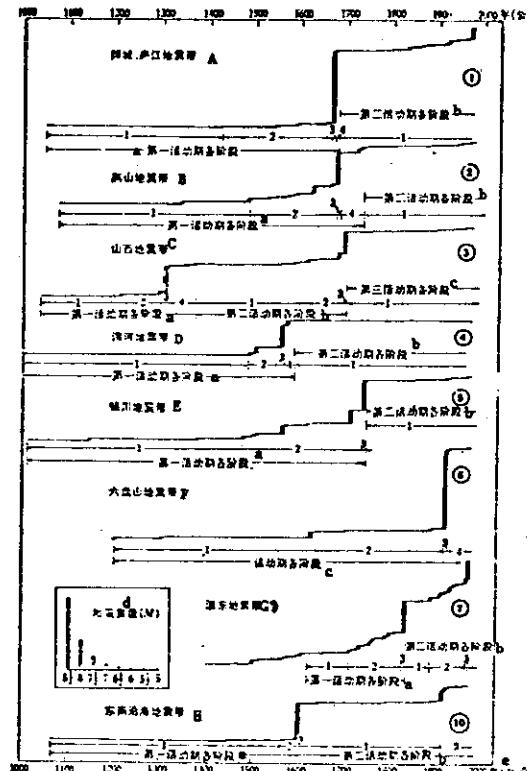
- $\beta/t_0 \rightarrow$ (decay const. of instrument)
(characteristic time of
mortality function)
- $M_0/M_0 \rightarrow$ unit moment
- $K (m/s) \rightarrow$ threshold moment

Input parameters \rightarrow (as above)

$$\begin{aligned} \beta/t_0 &= 4 \times 10^4 & (M_0 \text{ const}) \\ M_0/d_0 &\approx 3 \\ K &= 10^{-2} \end{aligned}$$

Input assumption of pairwise power law mortality (K_{pair})

- Weight
1. magnitude-frequency law
 2. Omori (1892) law of aftershocks
 3. Magnitude-length law
 4. Earthquakes non-stationary process
 5. Earthquakes may occur in supercluster where length exceed size spans of most catalogues
* China, Middle East, seismic catalogues

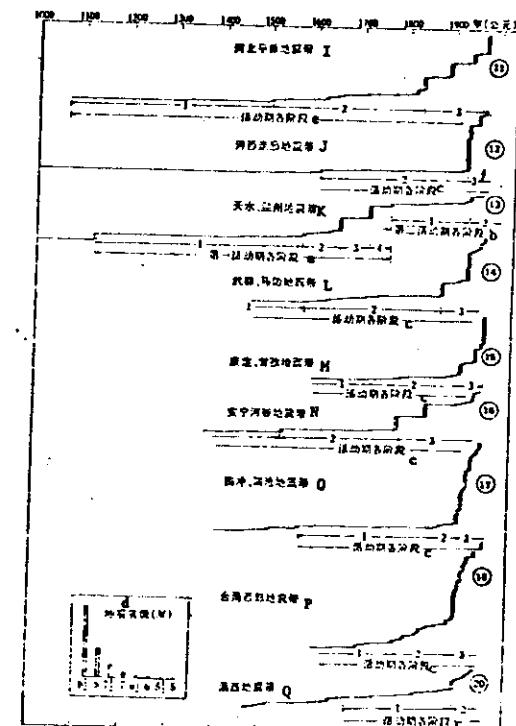


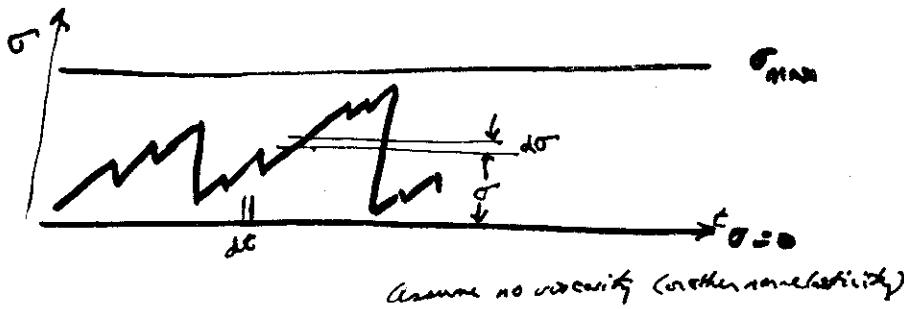
strain release curves for the seismic zones of China.

- a. Stages of the first seismic period.
- b. Stages of the second seismic period.
- c. Stages of the third seismic period.
- d. Magnitude.
- e. Year (AD).

- A. Tancheng-Lijiang seismic zone.
- B. Yanshan seismic zone.
- C. Shansi seismic zone.
- D. Weihe seismic zone.
- E. Yinchan seismic zone.
- F. Liupanshan seismic zone.
- G. Eastern Yunnan seismic zone.
- H. Southeast coast seismic zone.

- I. Hopei plain seismic zone.
- J. Hosi corridor seismic zone.
- K. Tianshu-Lanzhou seismic zone.
- L. Wudu-Mabian seismic zone.
- M. Kangding-Ganzi seismic zone.
- N. Anning valley seismic zone.
- O. Tengchong-Langcang seismic zone.
- P. Western Taiwan seismic zone.
- Q. Western Yunnan seismic zone.





$P(\sigma, t)/d\sigma$ is prob. that stress (strain) is between σ and $\sigma + d\sigma$

$\lambda(\sigma) dt \dots$ - an eq. will occur in time int. between $t \approx t + dt$ if we are in the state σ .

$T(x, \sigma) d\sigma \dots$ (cont.) if we are in state $x \in \mathbb{R}$, eq. occurs, then it tends to interval $\sigma \in \sigma + d\sigma$.

$$\lambda(\sigma) P(\sigma, t) + \alpha \frac{\partial P(\sigma, t)}{\partial \sigma} + \beta \frac{\partial P(\sigma, t)}{\partial t} = \int_{-\infty}^{\sigma + \Delta\sigma} P(x, t) d(x) T(x, \sigma) dx$$

Normalization

$$\int_{-\infty}^{\sigma} P(x, \sigma) dx = 1$$

Other

$$P(\sigma) d\sigma = 1$$

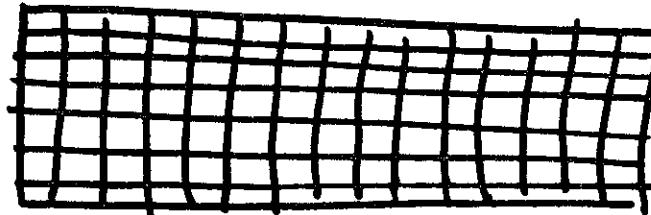
We find system is stable and stationary.

We expect interaction of time delays

$$\lambda(\sigma(t-t))$$

and produce clustering & (afterward?)

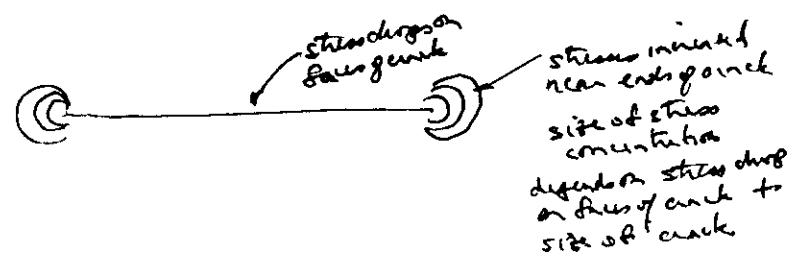
Model of fracture:

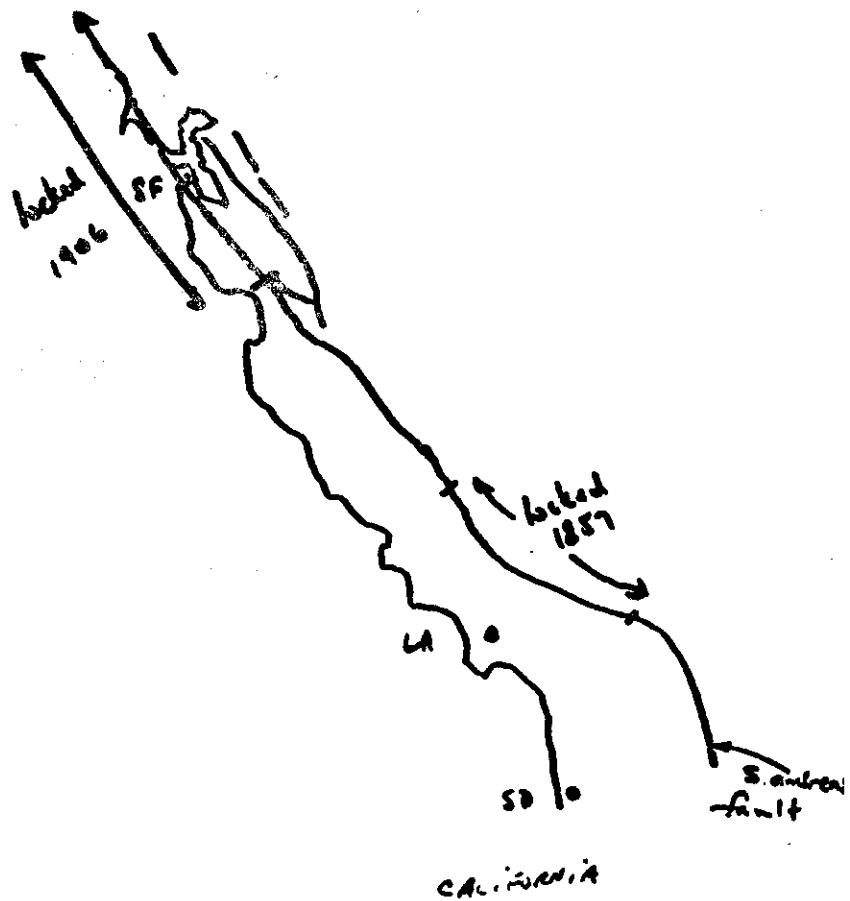


1. at each point (aspirity, barrier), we have a critical threshold of stress B (breaking strength) $B(x, y)$
2. To each point we apply an external force $F(x, y)$
3. Between events, $F(x, y)$ increases linearly with time (plate technique).
4. When fracture occurs crack grows quasi-statically.

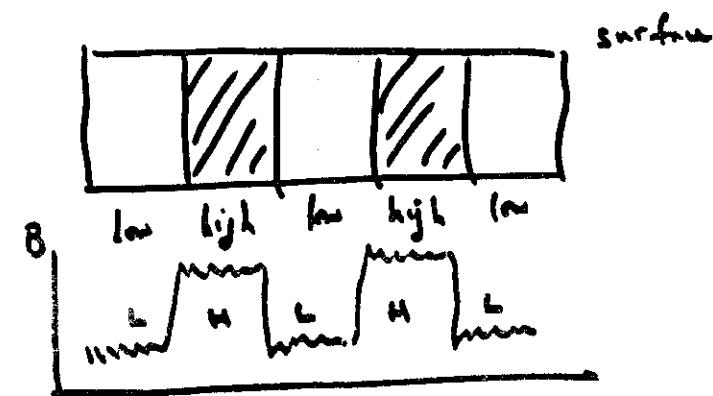


- 1. crack initiates some time $T(x, y)$ after $B = F$ at some point
- 2. Nearest neighbor search is made to see if redistributed stress, added to local value of F exceeds breaking of B etc.
- 3. When crack reaches its maximum size, stress in entire plane has been redistributed.

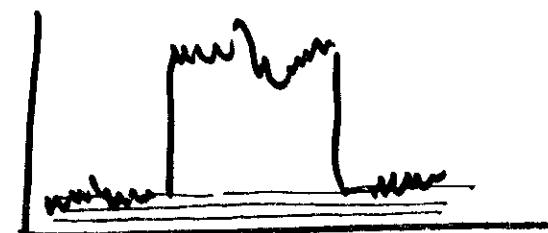




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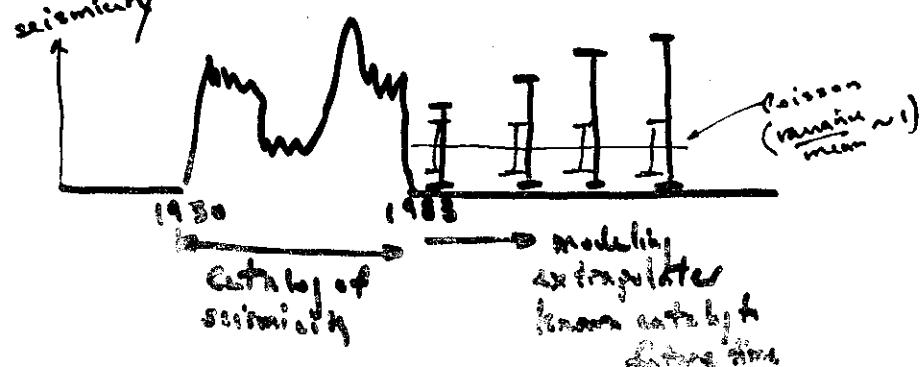
1. Time delays necessary to simulate magnitude-frequency law.
2. For most events, 99.5% of all events occur in L regions with correct mag. dep. law.
3. No events in L break completely across an L-H zone.
4. 0.5% of all events occur at an L-H boundary and tear completely across an H zone.
5. There are no small events in H.



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Seismic Risk Prediction

II-11



(if seismicity inter. indep.
 (nonindependent),
 then mean predictions are about
 the same as for Poisson model,
 but variances greater. In many
 other models (bursting process)
 variances may be much
 greater (e.g., 40 times average)

References

- Jawtooth Model (Holmogorov Filter)
 1. L. Knopoff, Rev. Geophys. & Space Phys., 9 (1971) 175.
 Banching Press
2. T.Y. Kagan & L. Knopoff, J. Geophys. Research,
 Crustal Seismicity
3. ~~Ma, ??~~ in Chinese Encyclopedia published by P.C.U.

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