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abdus salam

international centre for theoretical physics



H4.SMR/1150 - 12

Fifth Workshop on Non-Linear Dynamics and Earthquake Prediction

4 - 22 October 1999

The Simplicity of Generating Complexity

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Lecture I: The Simplicity of Generating Complexity

- Generating Power Laws
- Simple Examples of Complex Patterns
- Self-Organized Criticality (SOC)
- Earthquake Models of Complexity
- A CA Model Using Permeability as a Toggle Switch
- Application of the CA Model to Dehydration Reactions

Lecture II: Contemporary Models of the Earthquake Process

- Dislocations and Brittle Faulting
- Stress Transfer Models (e.g. NAFZ in Turkey)
- Coupling Fluid Flow Model to Large Scale Tectonics
- The Behavior of a 3-D Fluid-Controlled Earthquake Model

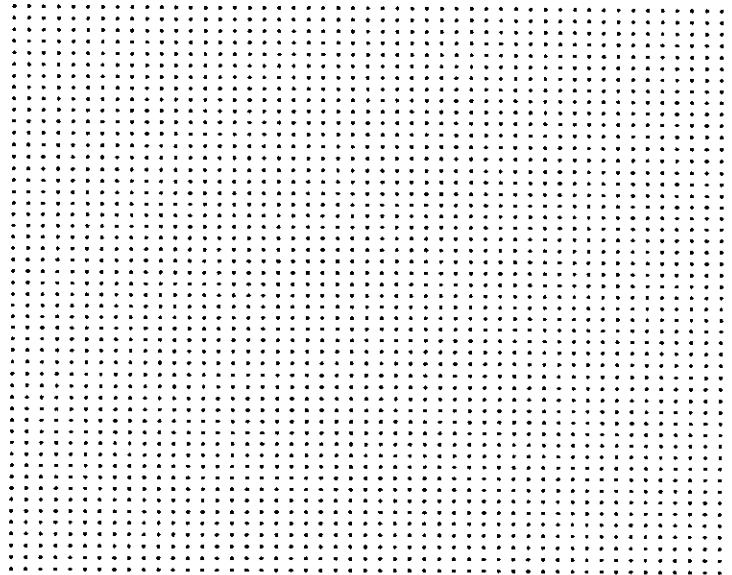
Lecture III: Properties of Large Ruptures and Surface Deformation Field

- Review of Some Rupture Models
- Complex Slip of Earthquakes
- General Strain Fields Around Elastic Dislocations
- The Properties of Large Ruptures (e.g. slip, rise time, moment release)
- Comparisons with Earhtquake Catalogs
- The Surface Deformation Field

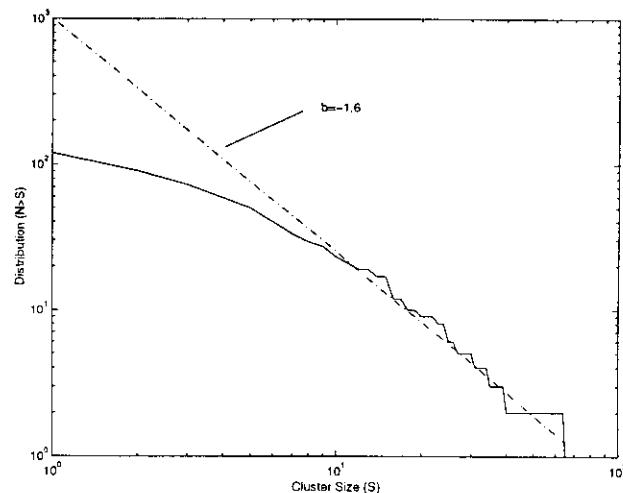
Objective: Build a model that includes:

- Complexity (Implied by Gutenberg-Richter)
 1. How to generate complexity
 2. Self-organized criticality (SOC)
 3. Molecular dynamics models
- Long-range stress interactions (Elasticity Theory)
 1. Brittle faulting as elastic dislocations
 2. Stress transfer models (earthquake triggering)
- Observation that faults are weak (In General)
 1. Stress measurements
 2. Heat flow anomalies
 3. Role of high pore pressures

Simple Complexity: The Dot Game



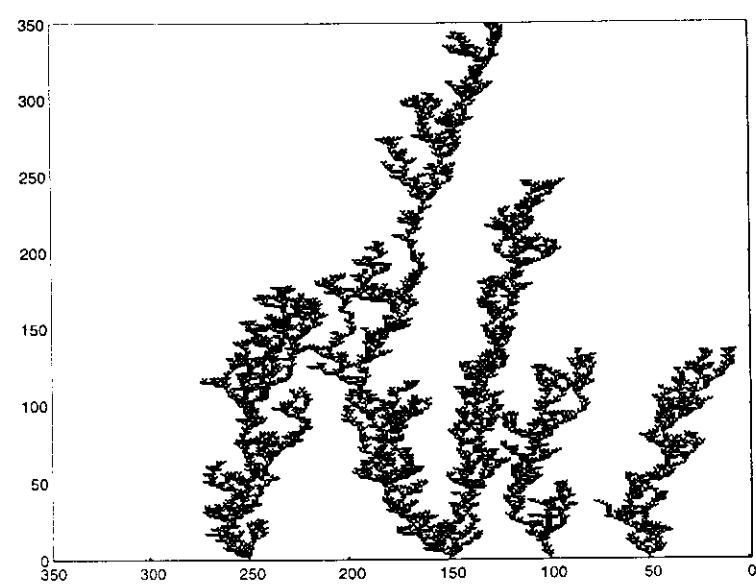
Connect the dots randomly until a box is created. If a box is enclosed, draw another line.
Continue turn until no more boxes can be enclosed. At the end of the game, count the
size of each box cluster created during one turn



Cluster size statistics. At the end of the game, count the size of each box cluster created
during one turn



Growth in an *A. oryzae* colony in a nutrient poor medium. From Matsuura and Miyazima (1993).



Simple Matlab program of positive feedback

```
colormap('gray')
V=[-1 0];
N=500;    matrix size
x=rand(N);  generate random number
y=x;
y(1:50:N)=-1;  perturb the boundary with a negative number
for k=1:30;  loop to search for connectivity to -1.
for j=2:N-1;
for i=2:N-1;
while y(i,j)>=0.95  & (y(i,j-1)==-1 | y(i+1,j)==-1 | y(i-1,j)==-1)
y(i,j)=-1;
y(i+1,j+1)=-1;
y(i-1,j+1)=-1;
y(i,j+1)=y(i,j+1)+.1;
y(i+1,j)=y(i+1,j)+.1;
y(i-1,j)=y(i-1,j)+.1;
end;
end;
end;
step forward in time
y(y>=0)=y(y>=0)+.01; step forward in time
h=pcolor(y);caxis(V);shading flat;pause(.5); plot results
end;
```

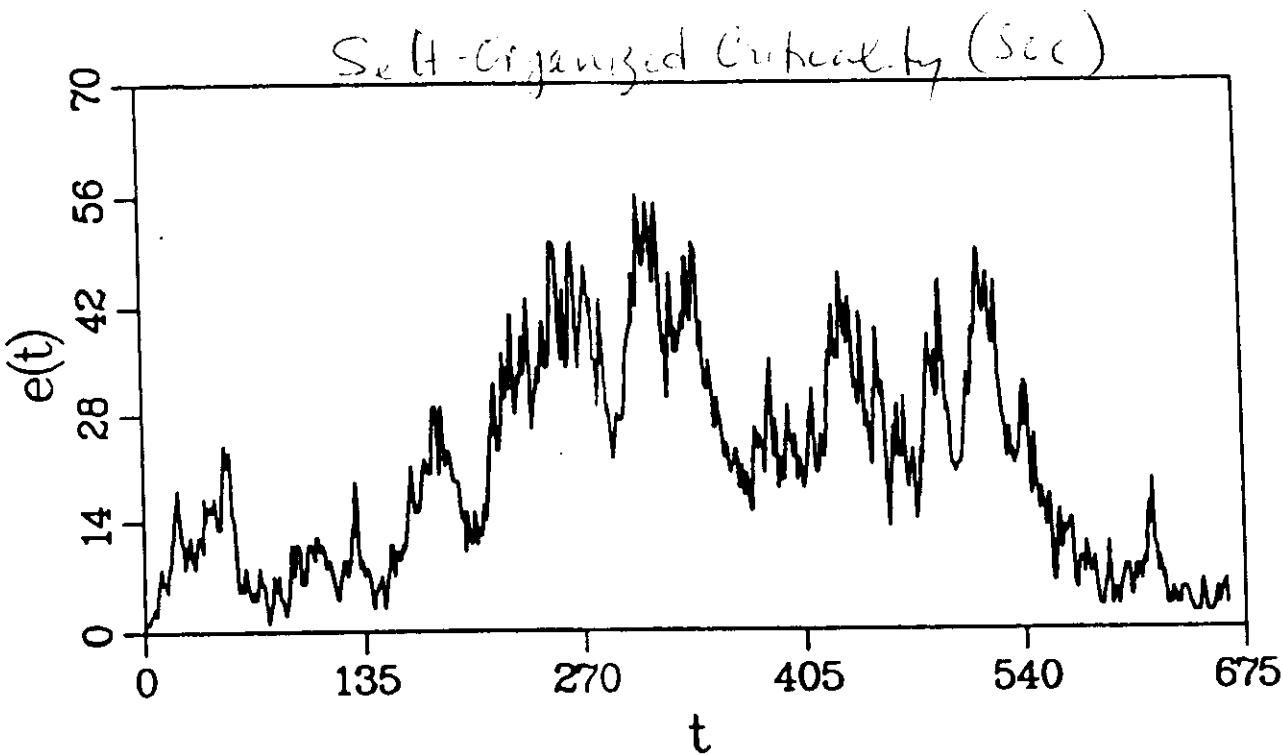
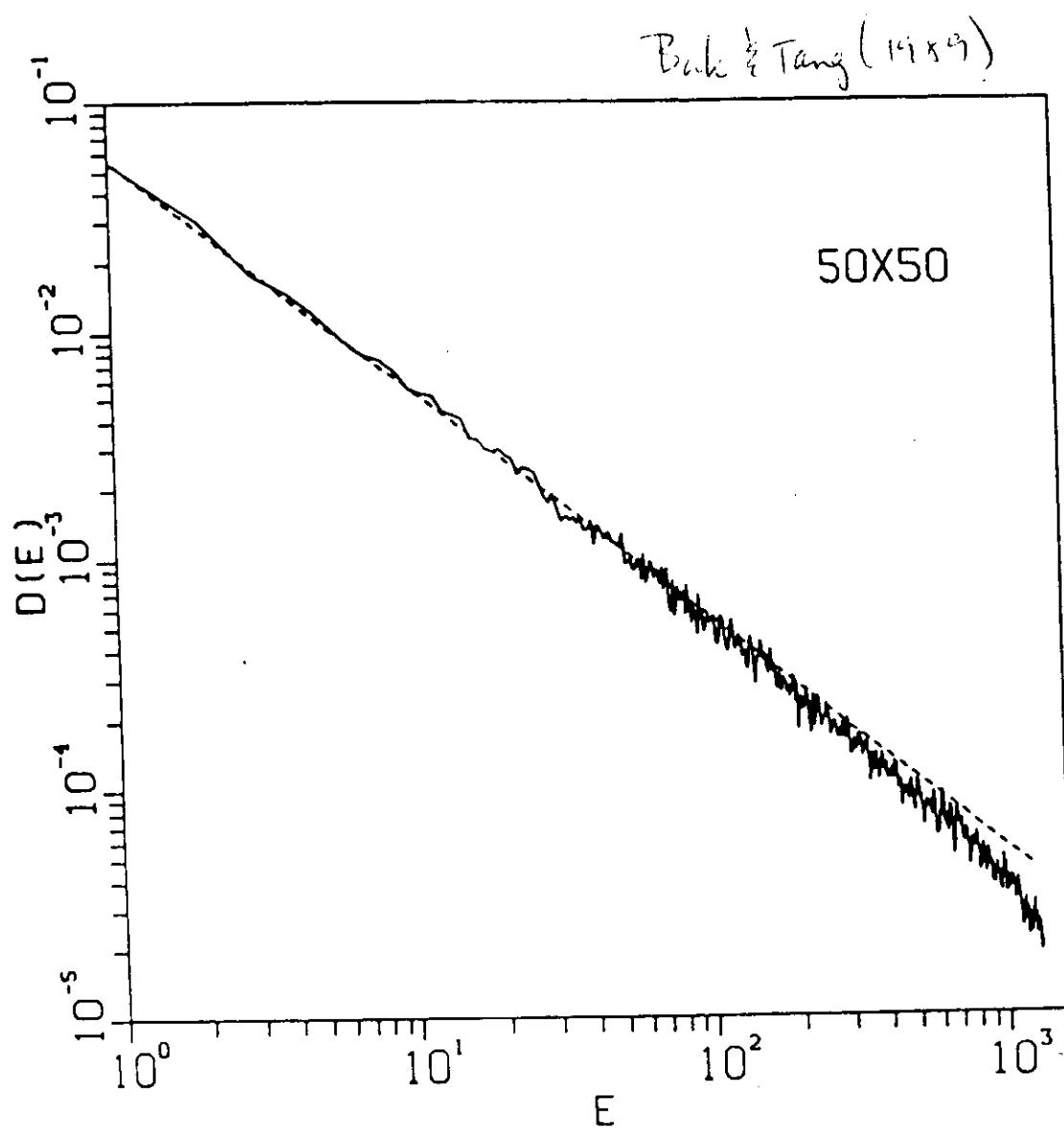
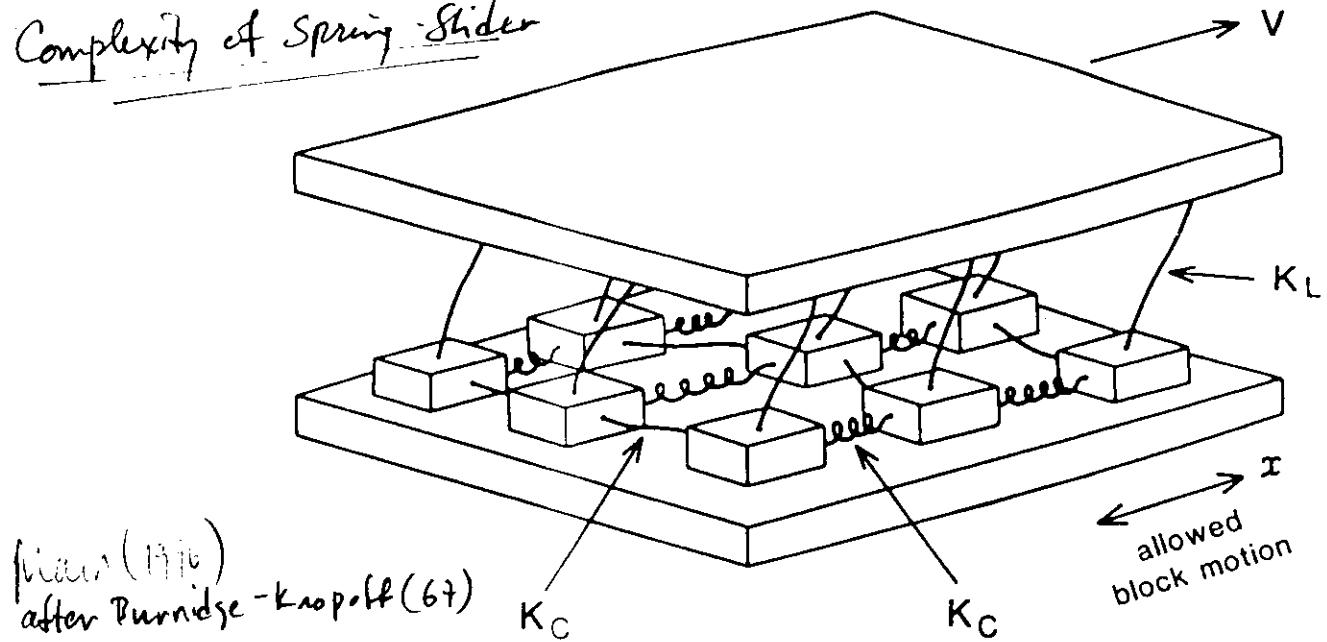


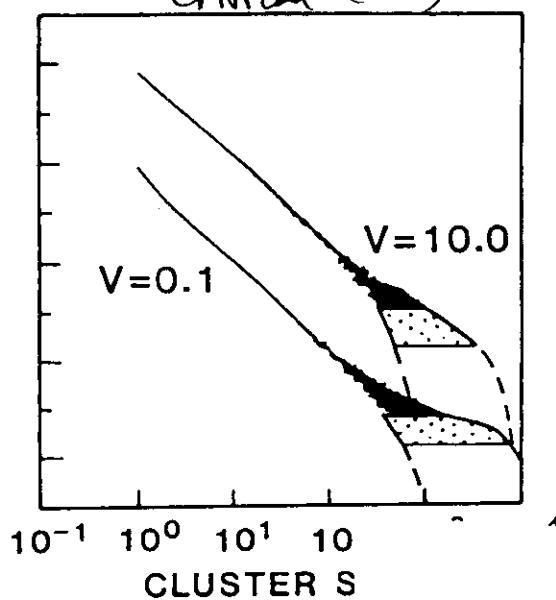
Fig. 1. Energy release versus time during a typical earthquake.



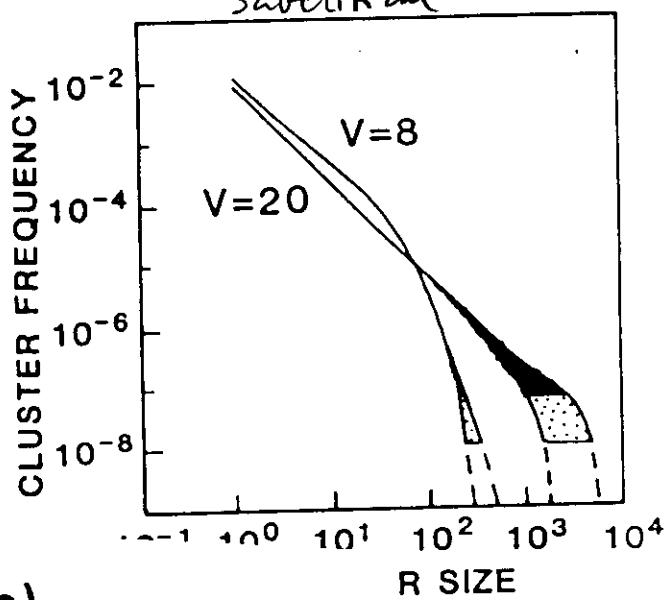
Complexity of Spring Slider



(b)
Critical (soc)

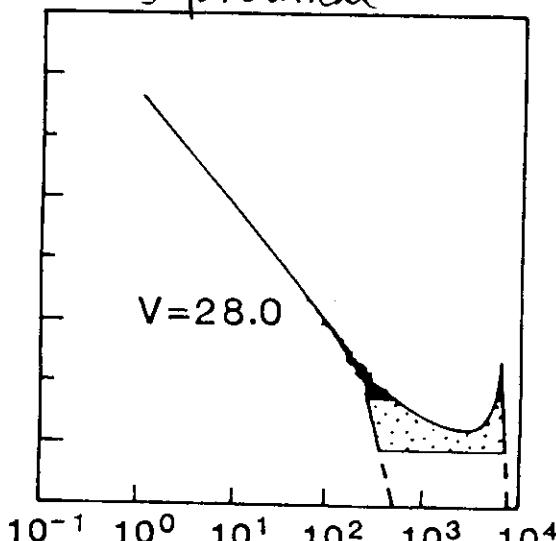


(a)
Subcritical



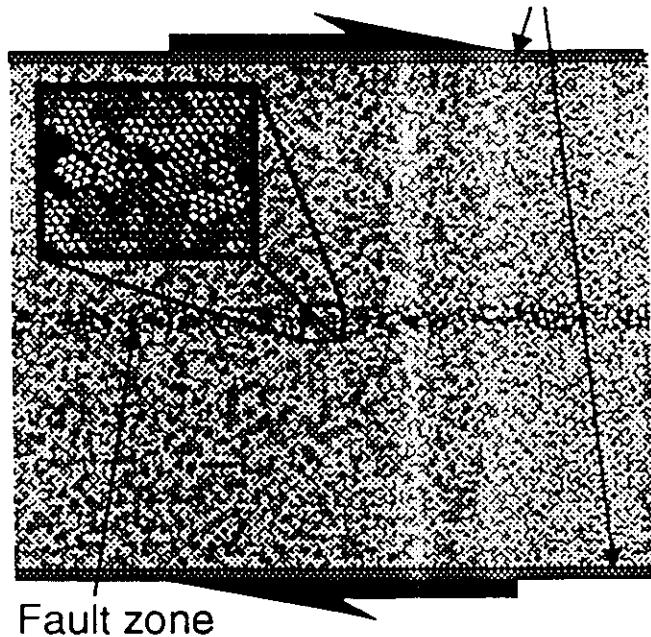
(c)

Supercritical

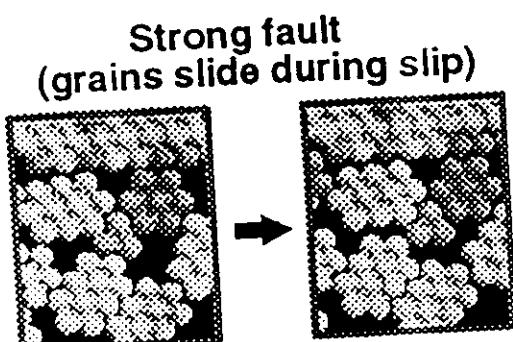
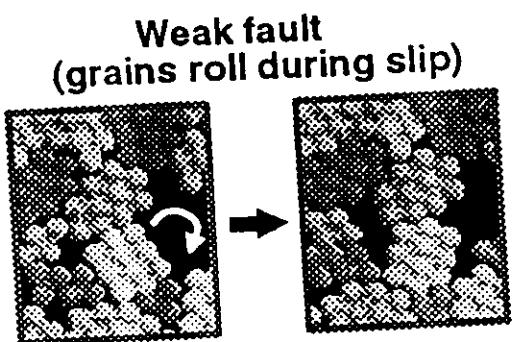


from Rundle & Klein (93)

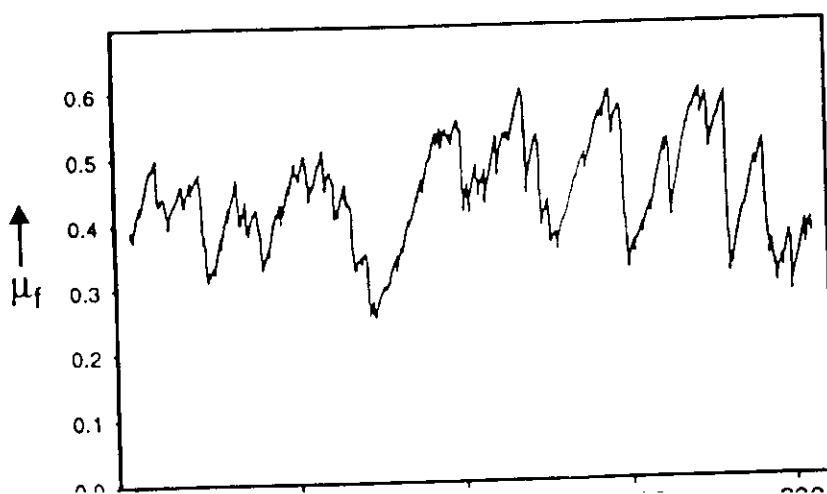
Rigid driving plates



Lattice-Solid Model
of Fault zones

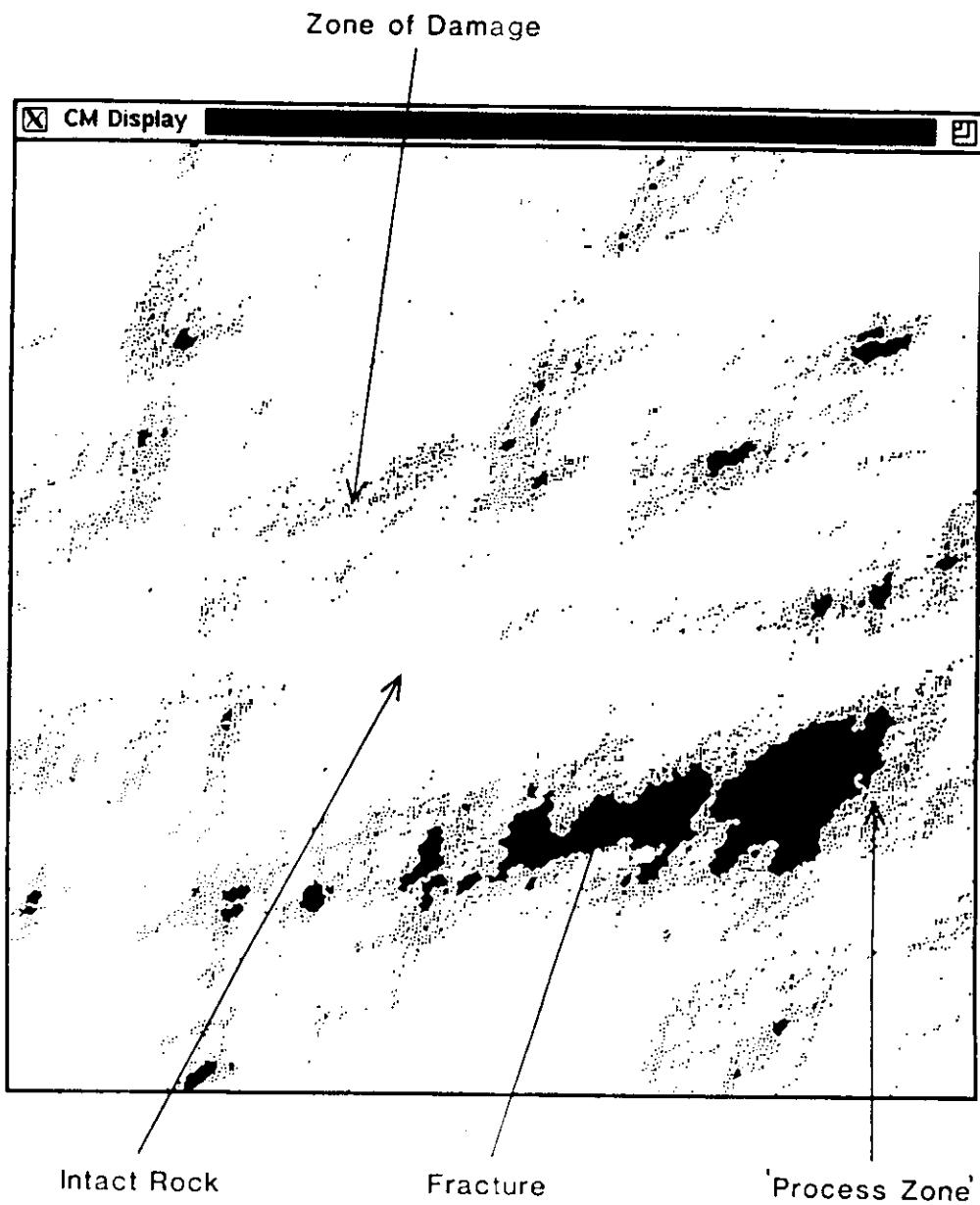


Mora & Plante (98)



Cellular Automata of Damage

Henderson et al (94)



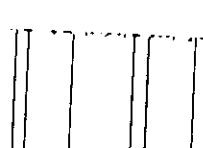
wolfram(86)



rule 90 (01011010)



rule 94 (01011110)



rule 104 (01101000)



rule 105 (01101001)



rule 106 (01101010)



rule 108 (01101100)



rule 110 (01101110)



rule 122 (01110101)



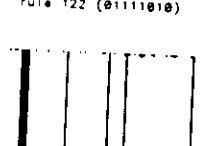
rule 126 (01111110)



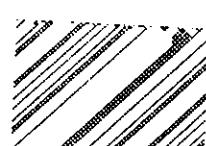
rule 128 (10000000)



rule 130 (10000010)



rule 132 (10000100)



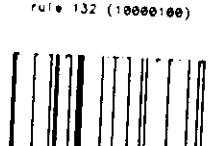
rule 134 (10000110)



rule 136 (10001000)



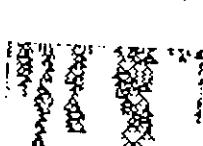
rule 138 (10001010)



rule 140 (10001100)



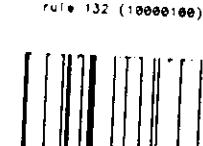
rule 142 (10001110)



rule 146 (10010010)



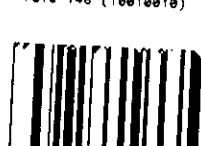
rule 150 (10010110)



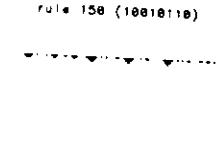
rule 152 (10011000)



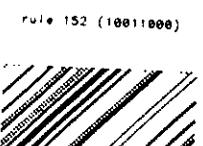
rule 154 (10011010)



rule 156 (10011100)



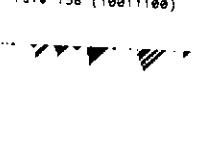
rule 158 (10010110)



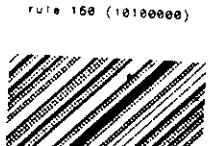
rule 160 (10100000)



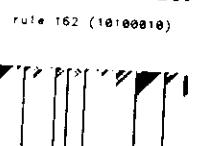
rule 164 (10100100)



rule 168 (10101000)



rule 170 (10101010)



rule 172 (10101100)

FAULT VALVE BEHAVIOR

SIBSON ('73, '92)

CRUSTAL STRESS, FAULTING & FLUID FLOW

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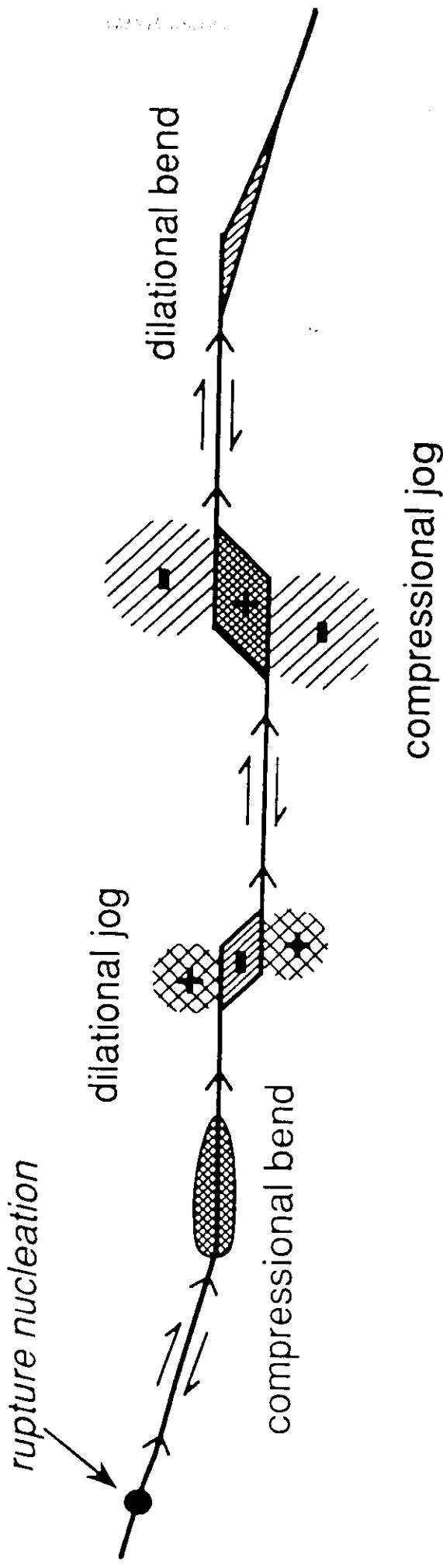
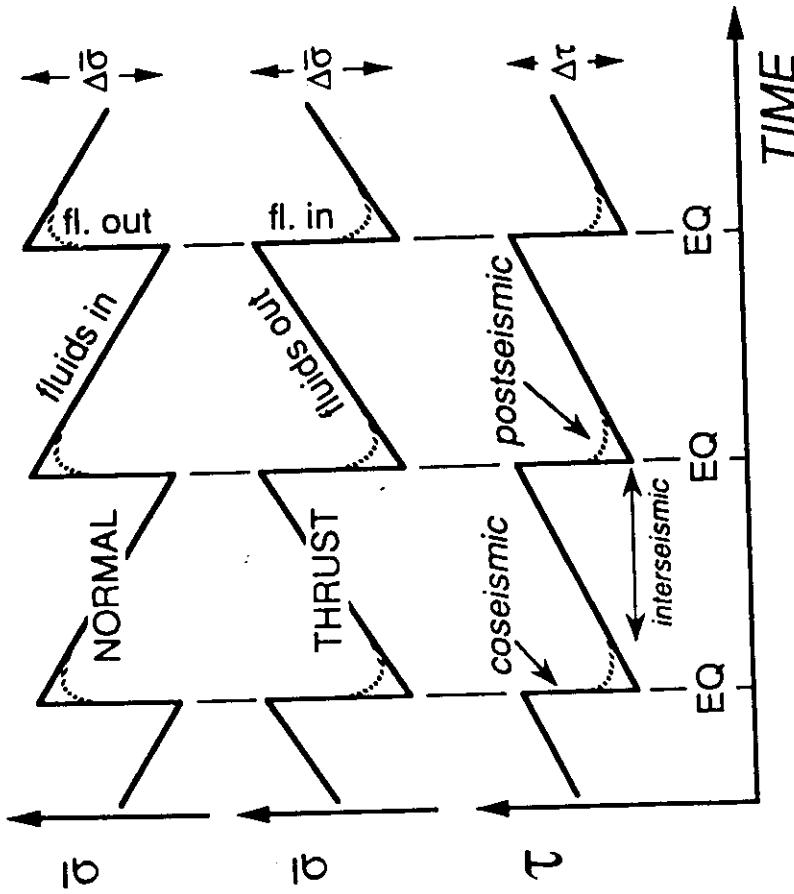
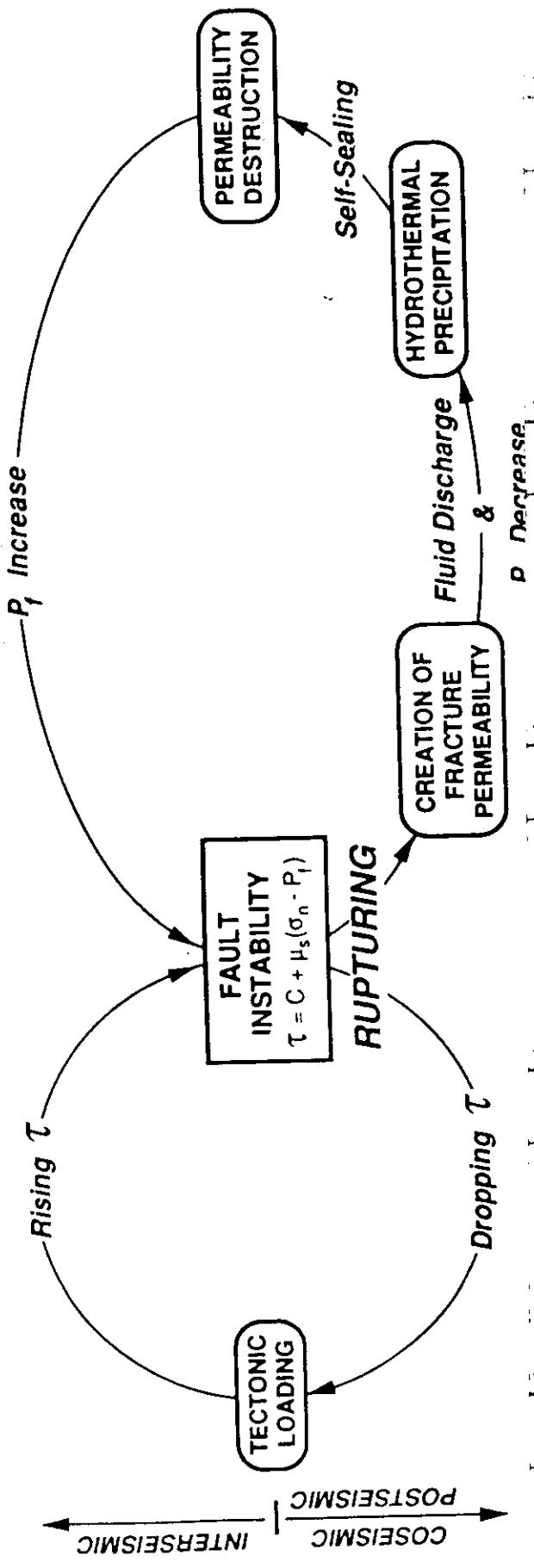


Fig. 5. Seismotectonic carbon of an irregular rupture trace (not to scale), showing areas of enhanced (+ and cross-hatched) and reduced (— and diagonal hachures) mean stress arising from a rupture propagating from left to right. Note that the response of isolated fault bends depends on the direction of rupture propagation. Diagram represents a map view of a strike-slip fault, or a cross-section through a dip-slip fault.

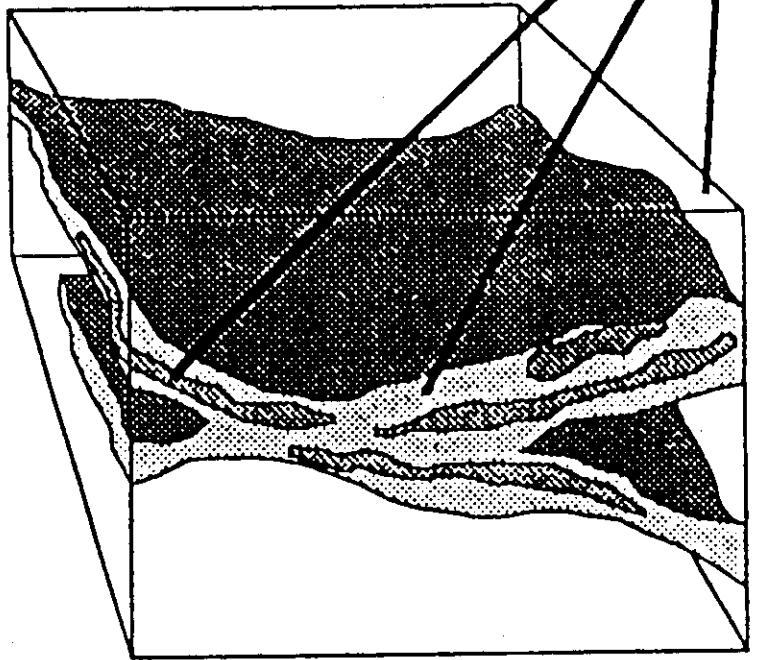


FAULT VALVE BEHAVIOR
SIBSON ('92)

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FAULT ZONE ARCHITECTURAL COMPONENTS



FACTORS CONTROLLING k

- Lithology
- Fault scale
- Fault type
- Deformation style & history
- Fluid chemistry
- P-T history
- Component percentage
- Component k
- Component anisotropy (magnitude & direction of k_{\max} & k_{\min})

Figure 1. Conceptual model of fault zone with protolith removed (after Chester and Logan, 1986; Smith et al., 1990). Ellipse represents relative magnitude and orientation of the bulk two-dimensional permeability (k) tensor that might be associated with each distinct architectural component of fault zone.

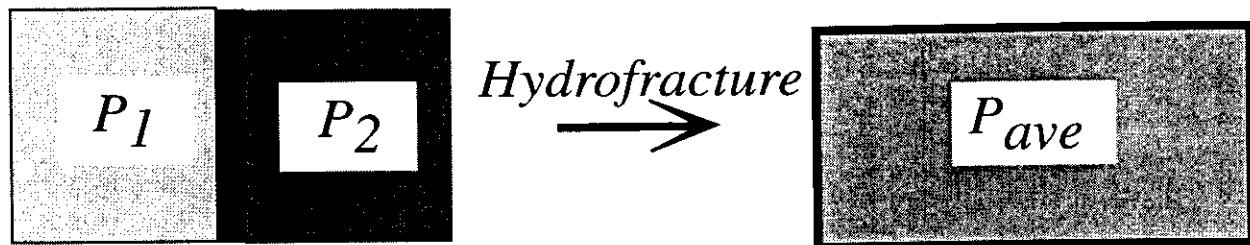
CANER et. al (1996)

CELLULAR AUTOMATA of CRUSTAL FLUID FLOW

Objective:

Answer the simple question:

"*What happens if two neighboring cells at different pore pressures suddenly communicate?*"

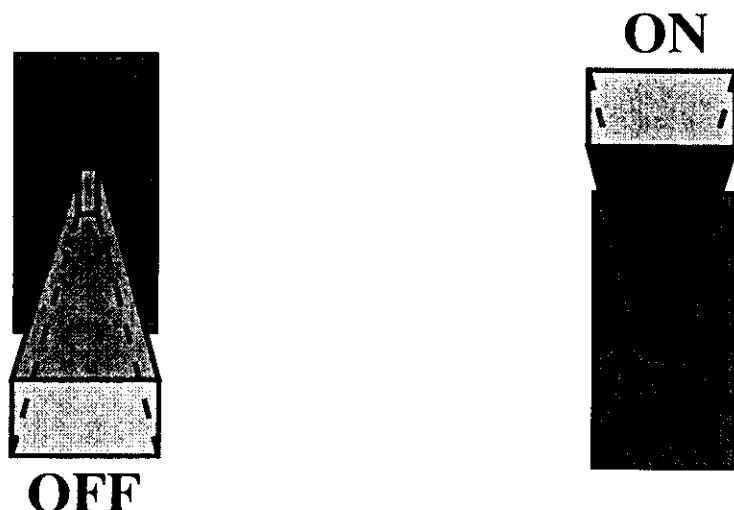


Approach:

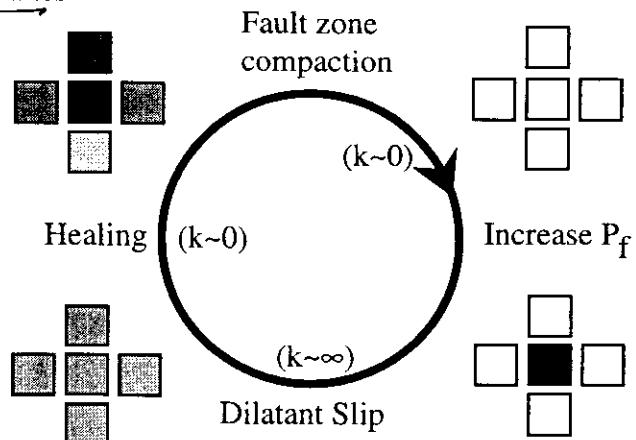
- Investigate a system where permeability is treated as a *toggle-switch*
- Apply the model to approximate fluid flow in fault zones, couple to *elastic dislocation* theory, and build a fluid-controlled *earthquake model*
- Apply the model to *dehydration reactions* and compare with laboratory experiments

INTERACTION RULE

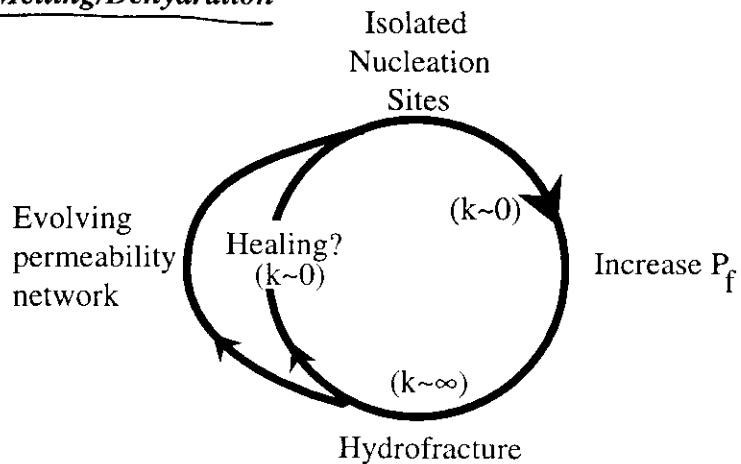
Permeability as a Toggle Switch



a) Earthquakes



b) Melting/Dehydration



Diffusion Equation with Source Term

$$\frac{\partial P_f}{\partial t} = \frac{1}{\phi(\beta_\phi + \beta_f)} \left[\frac{k}{v} \nabla^2 P_f - (\dot{\phi}_{plastic} - \dot{\Gamma}) \right] \quad (1)$$

where

- ϕ is porosity
- β_ϕ and β_f are the pore and fluid compressibility
- v is the viscosity
- ρ is the fluid density
- k is the intrinsic permeability of the matrix.
- $\dot{\phi}_{plastic} - \dot{\Gamma}$ is a source term

For an impermeable matrix ($k \sim 0$)

$$\frac{\partial P_f}{\partial t} |_{noflow} = \frac{(\dot{\Gamma} - \dot{\phi})_i}{\phi_i \beta_i} \quad (2)$$

At failure (hydrofracture), fluid pressures equilibrate with nearest neighbor cells. The equilibrium pressure (by conserving fluid mass and ignoring gravity, is:

$$\bar{P} = \frac{\sum_{i=1}^m (\phi \beta)_i P_i}{\sum_{i=1}^m (\phi \beta)_i} \quad (3)$$

where \bar{P} is the average pressures of affected cells.

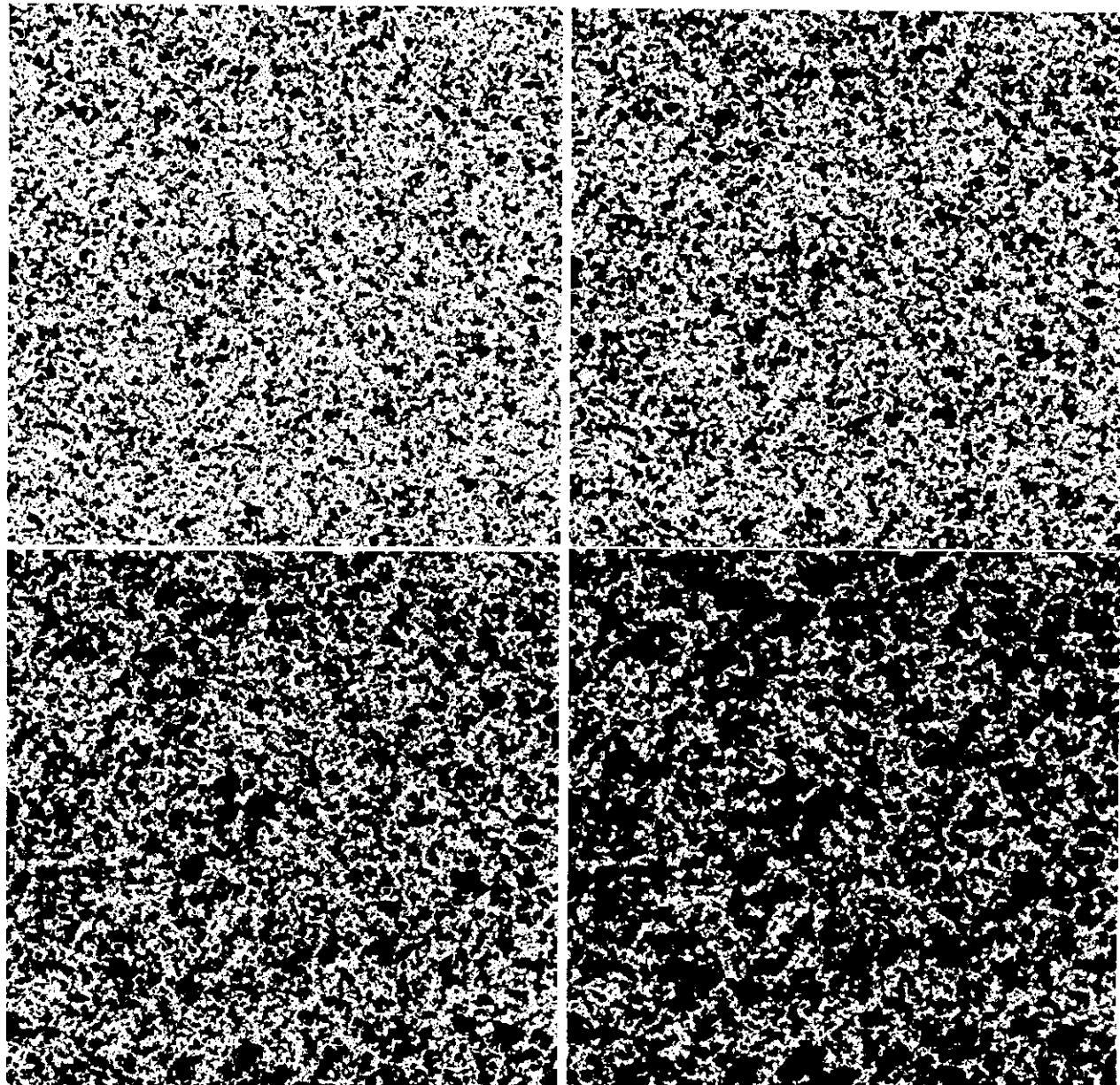
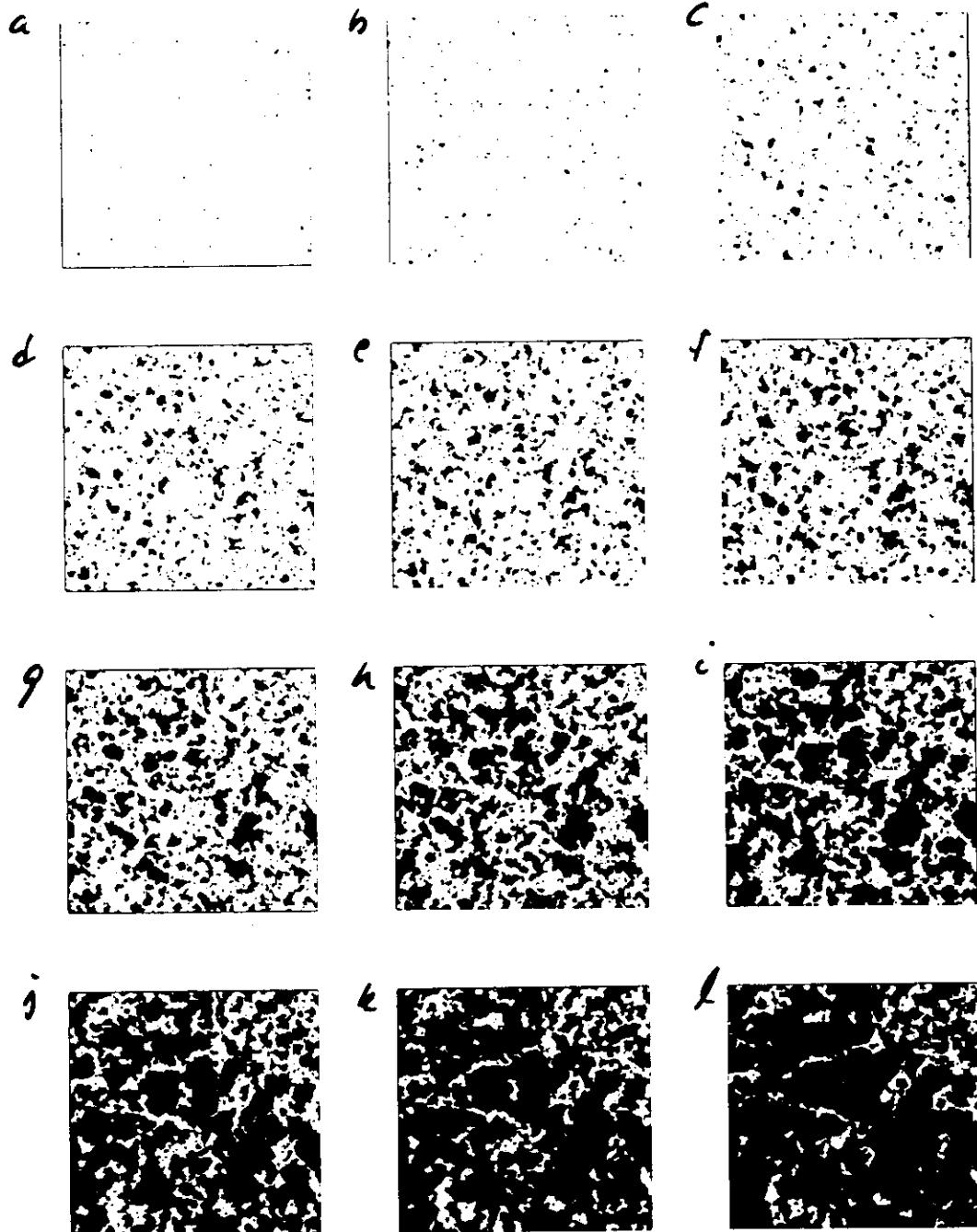


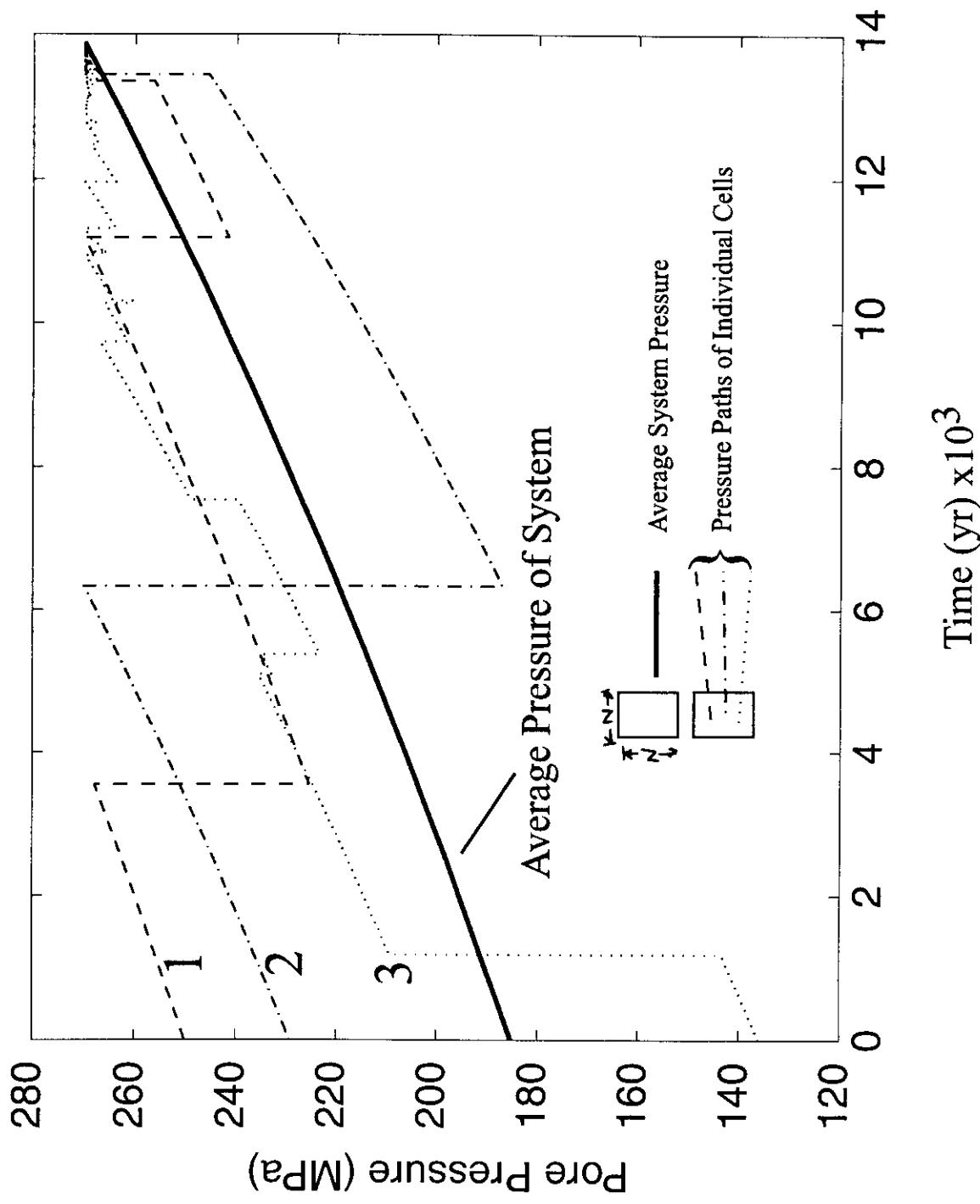
Figure 5: Evolution of high pore pressure connectivity from cellular automaton model

EVOLUTION OF INCIPIENT FAILURE

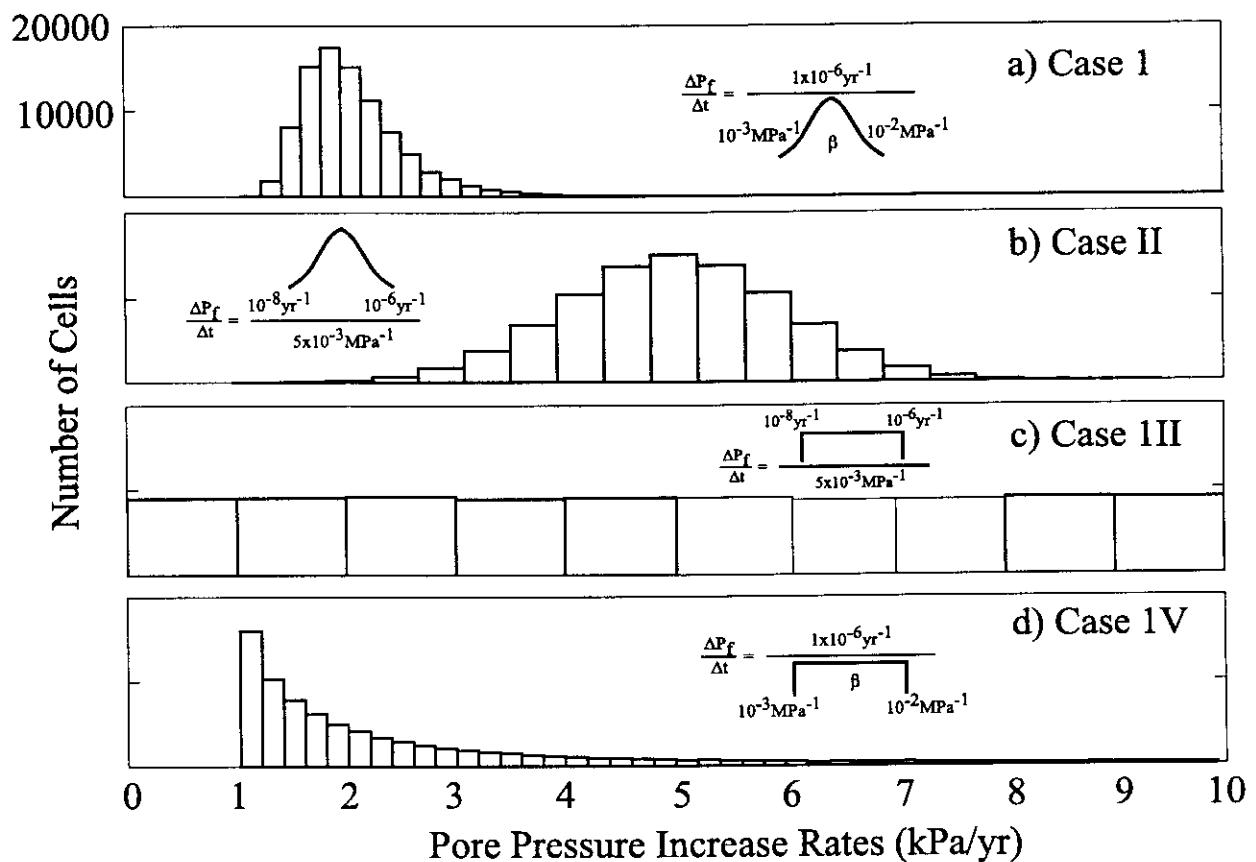


Black = Incipient Failure

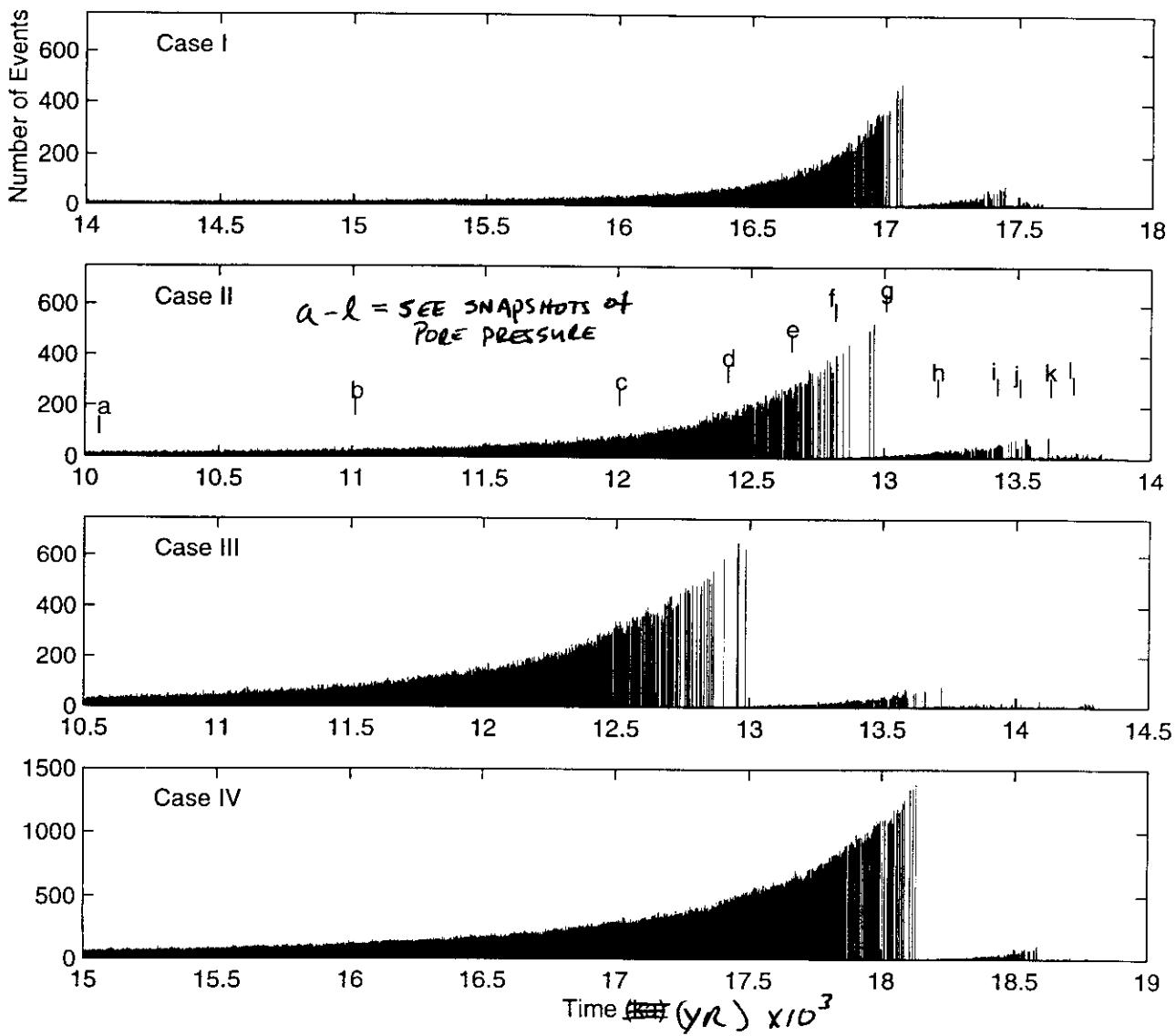
Pore pressure paths of arbitrary cells and system pressure



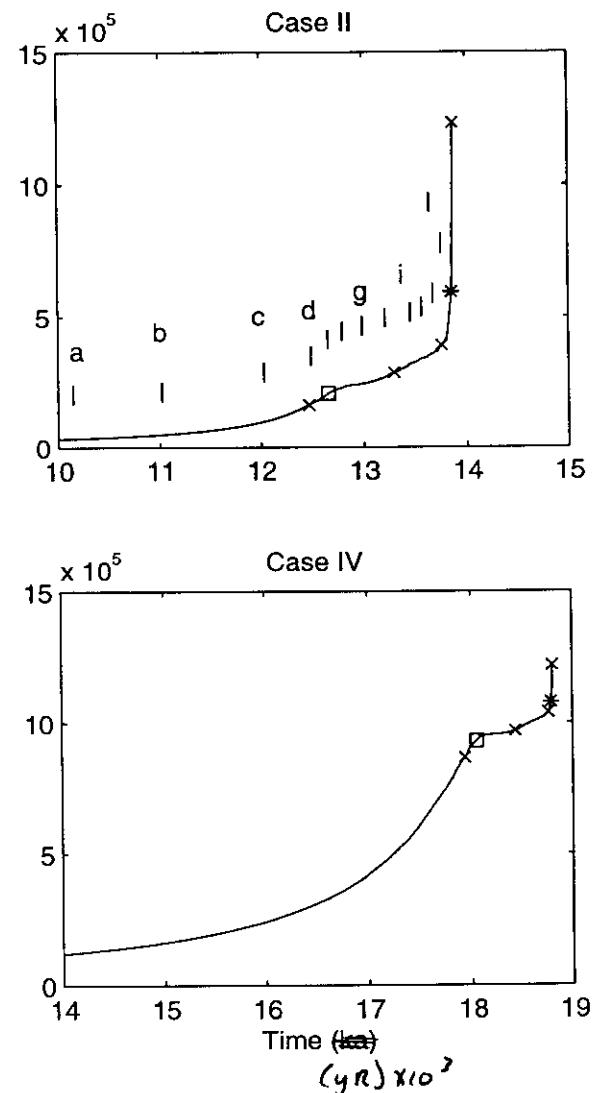
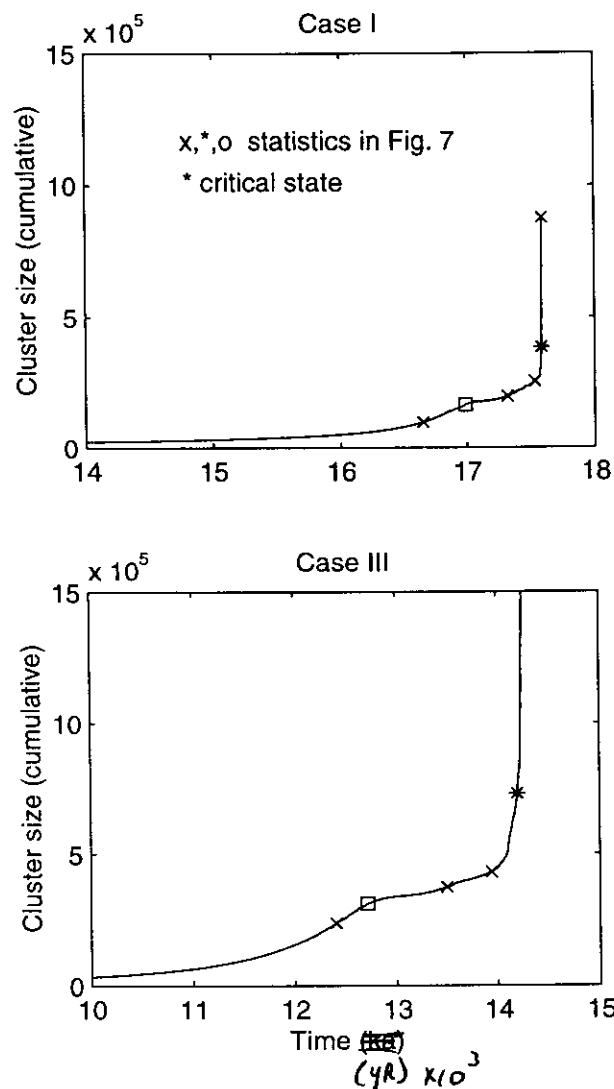
INPUT TO CA MODEL ; FOUR CASES of
PORE PRESSURE INCREASE RATE DISTRIBUTIONS

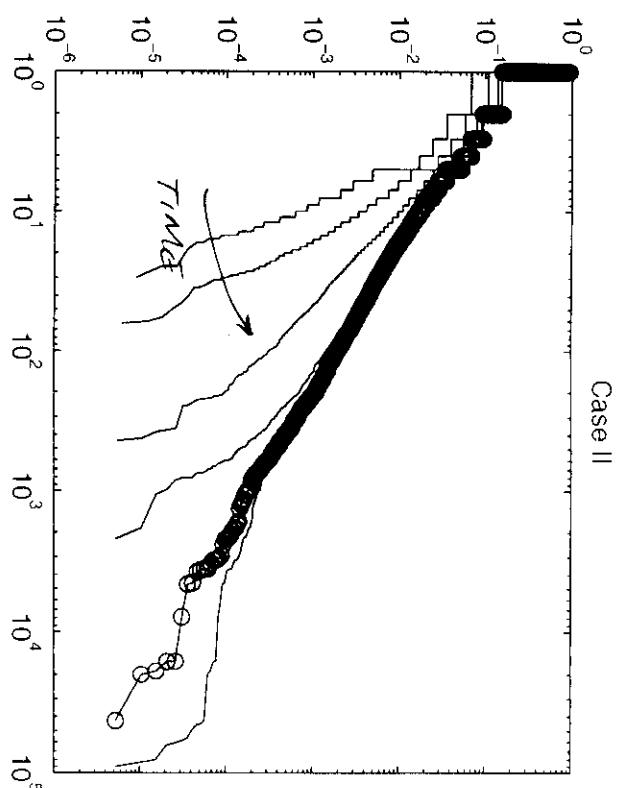
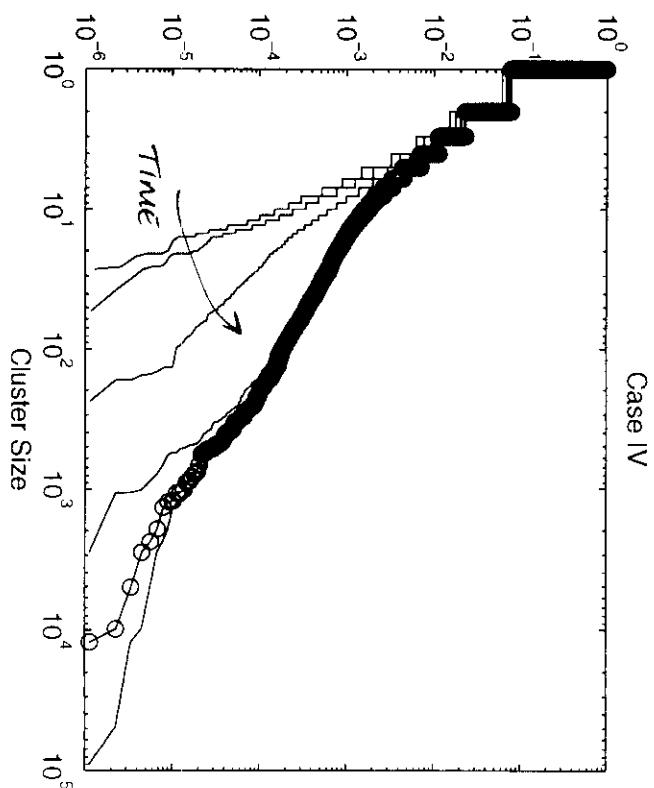
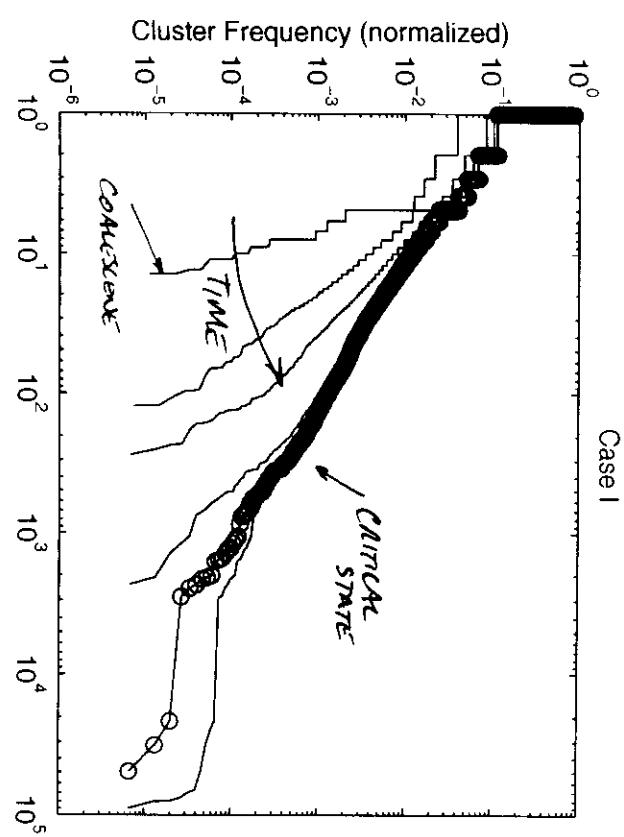
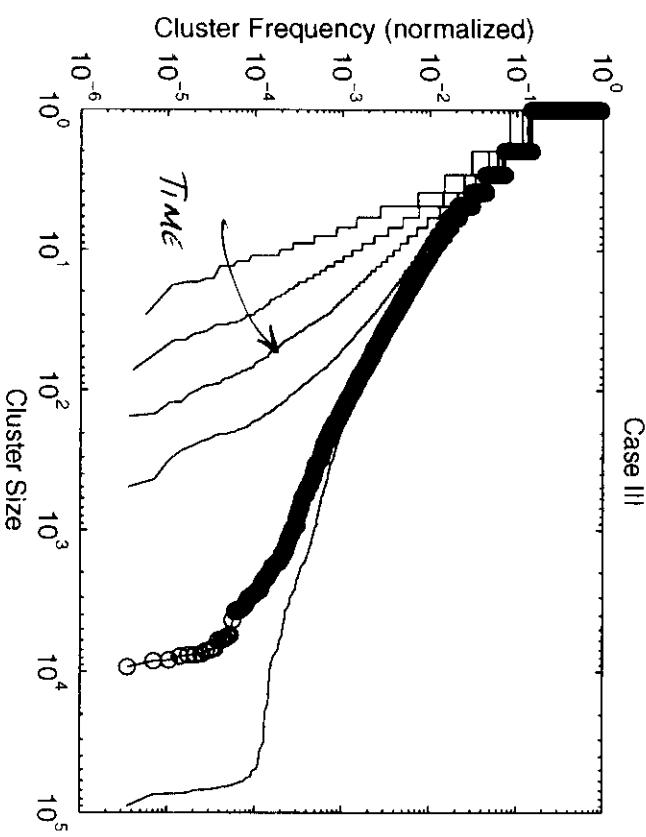


TIMELINE of Number of EVENTS

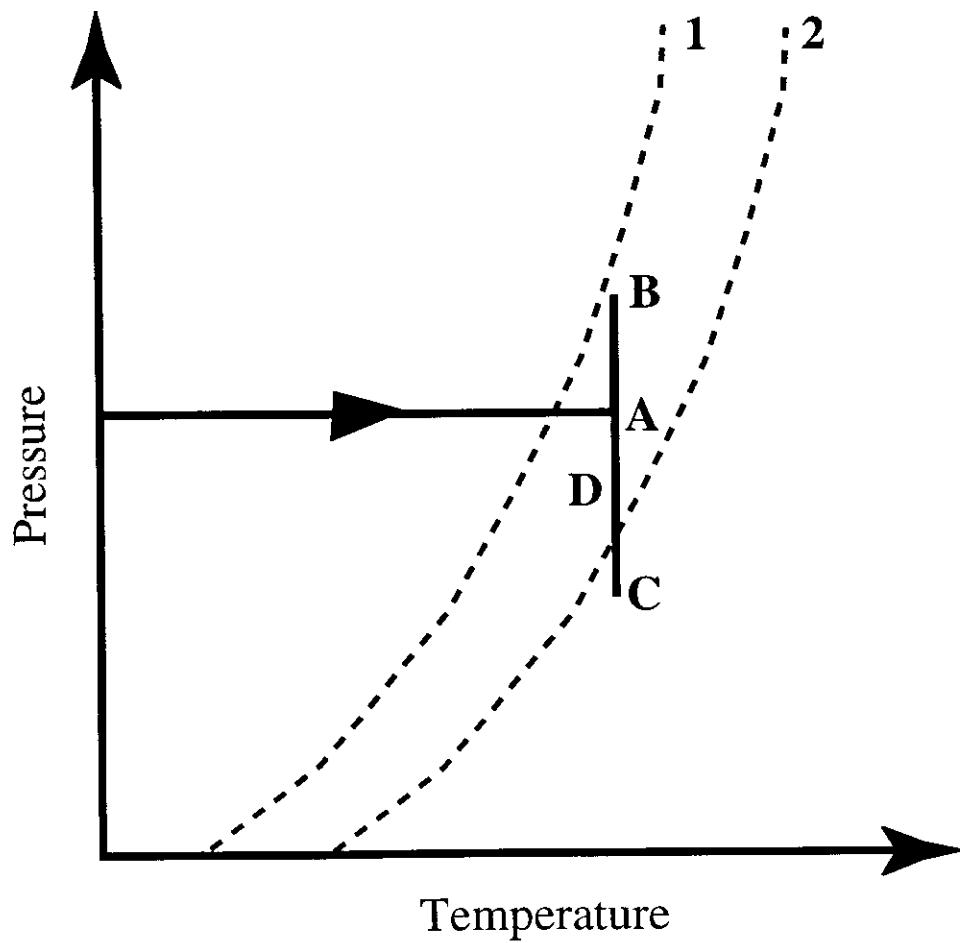
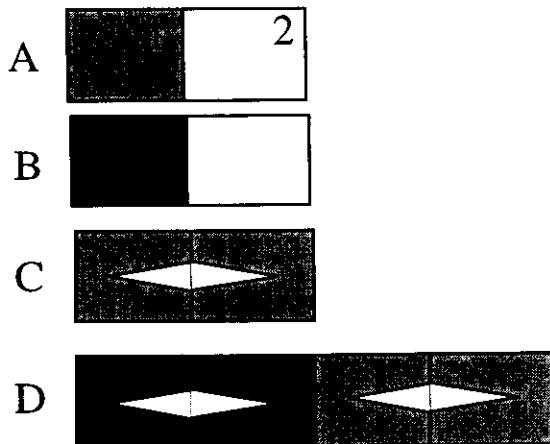


Evolution of CLUSTER SIZES



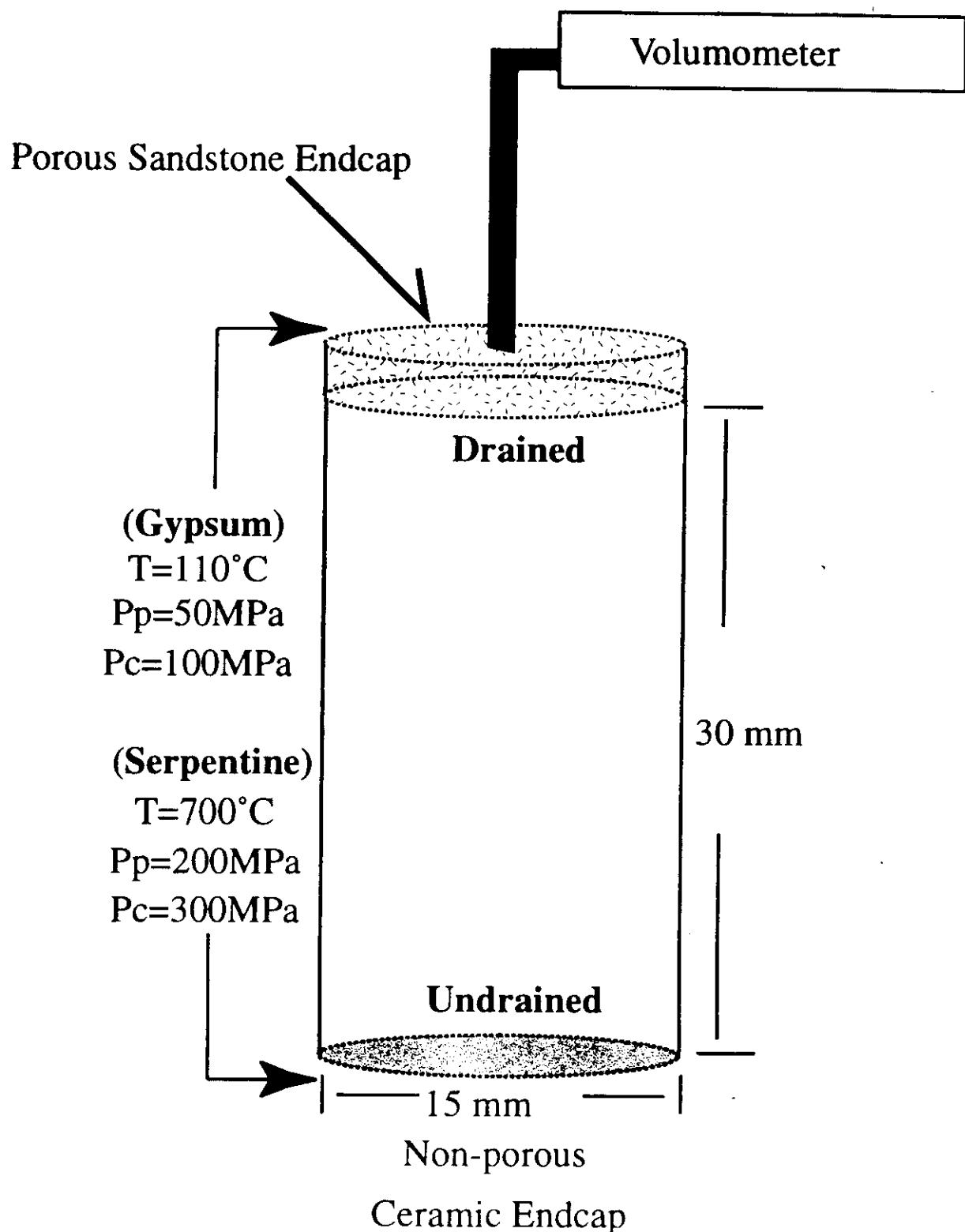


CONCEPTUAL MODEL of DEHYDRATION REACTIONS
USING TOGGLE-SWITCH PERMEABILITY ASSUMPTION

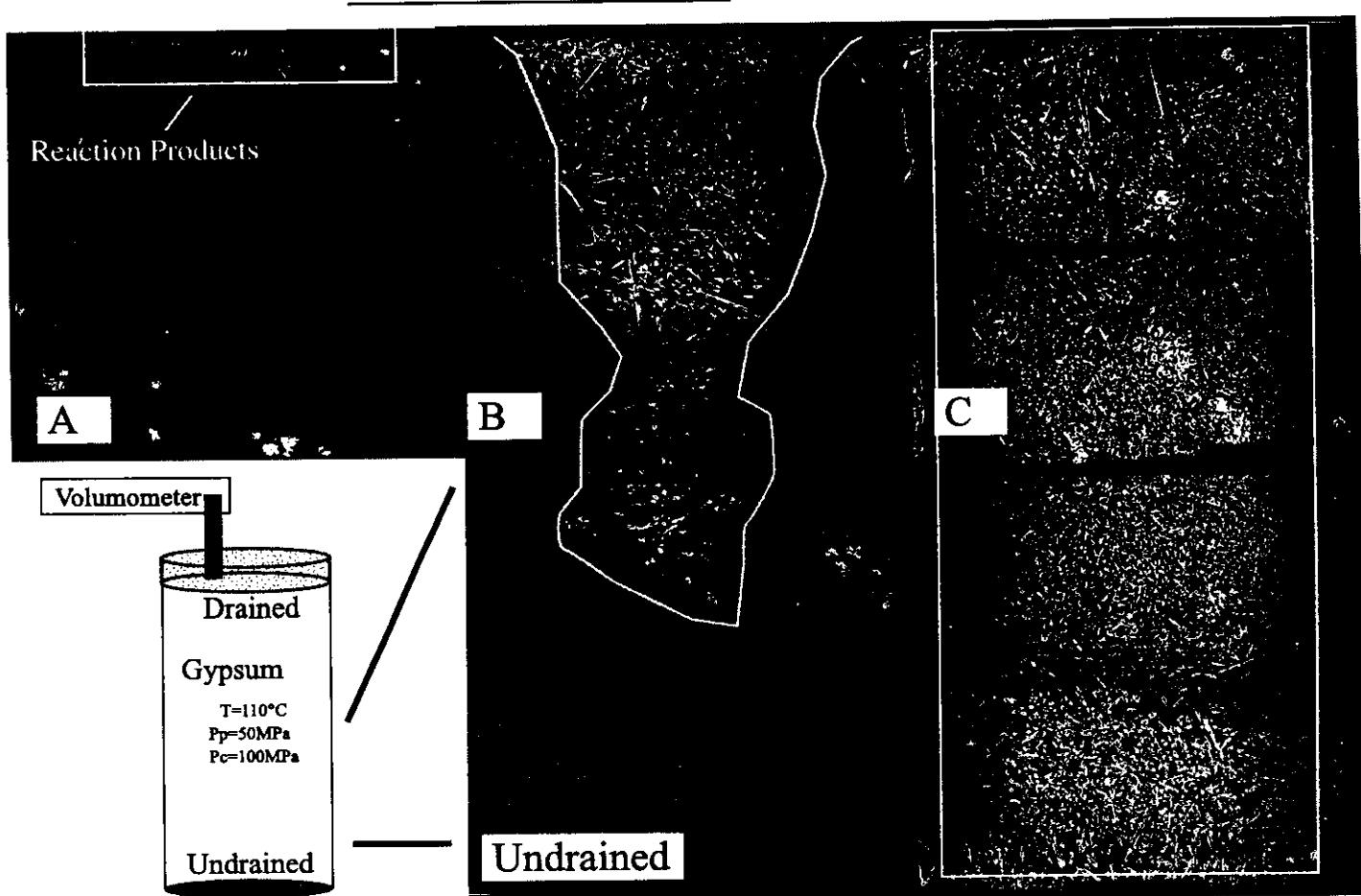


Miller, vanderzee, Olgard, Connolly ('99, submitted)

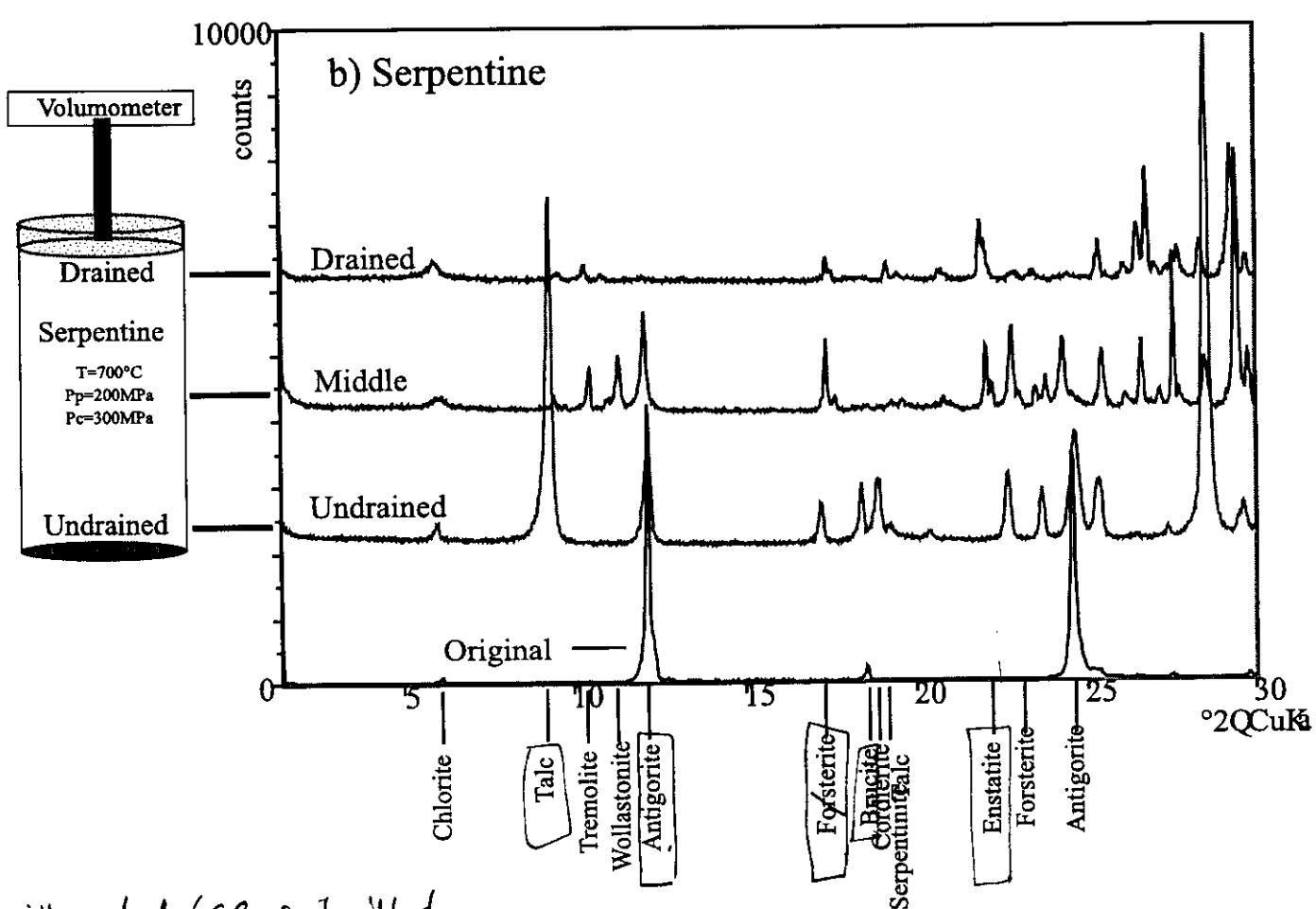
Experimental Setup



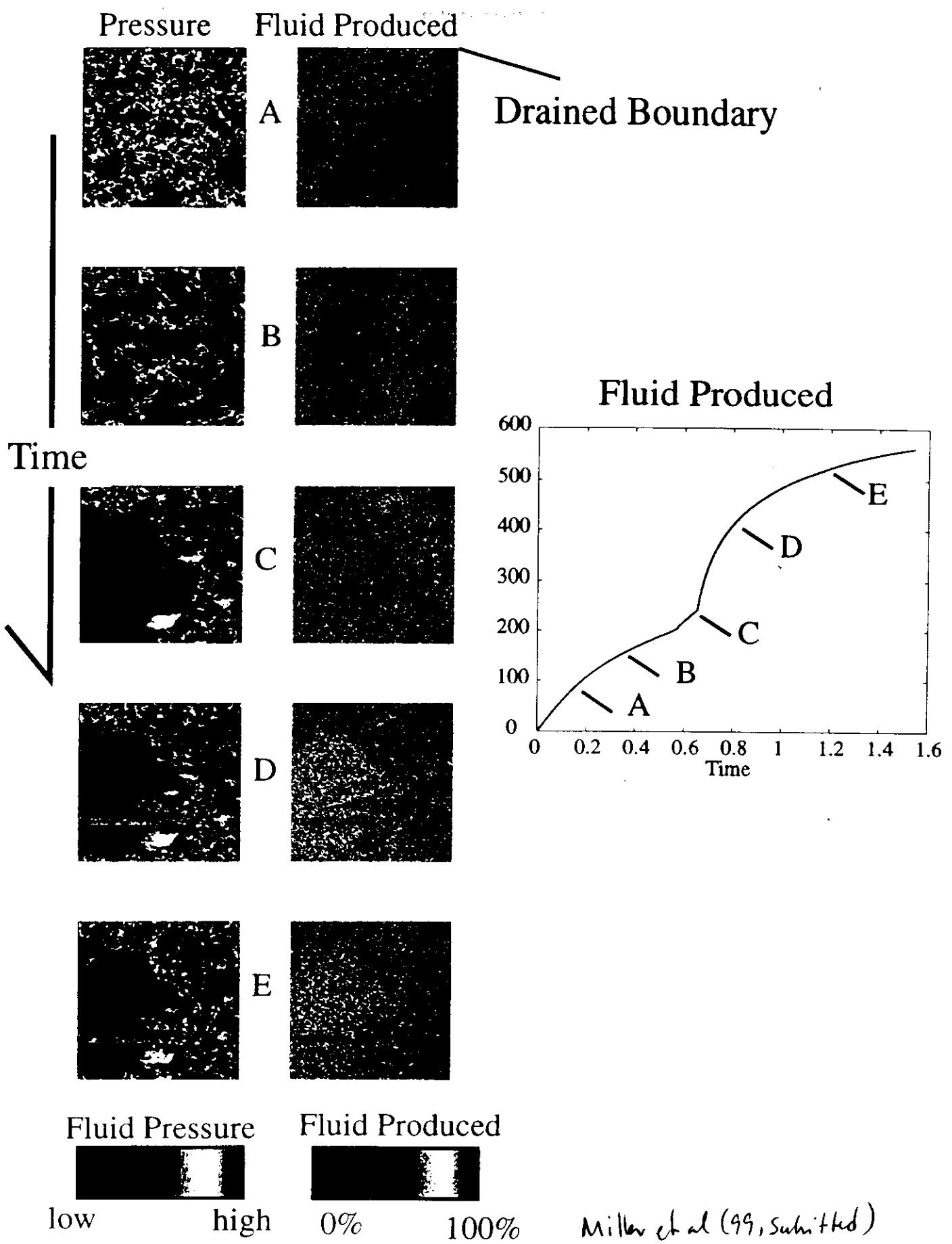
Miller et al. (95, submitted)



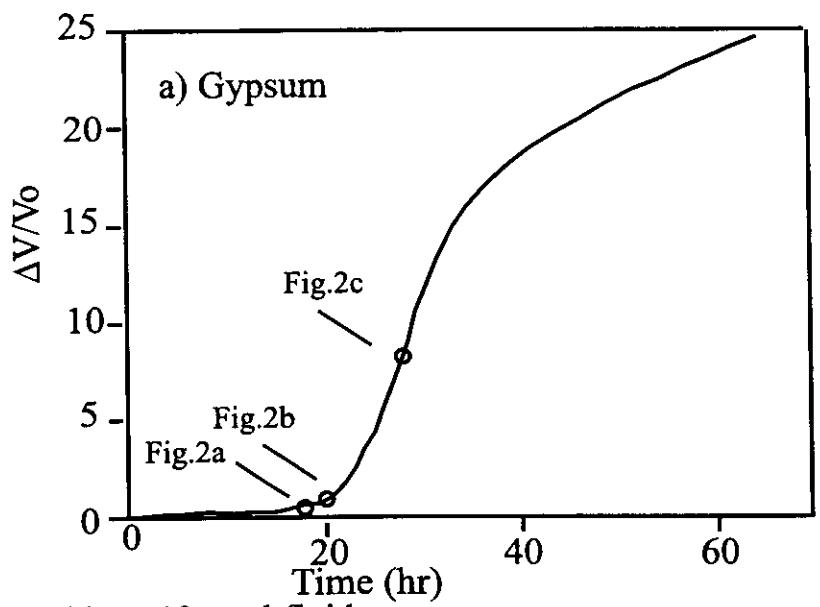
Van der Zee, (97)



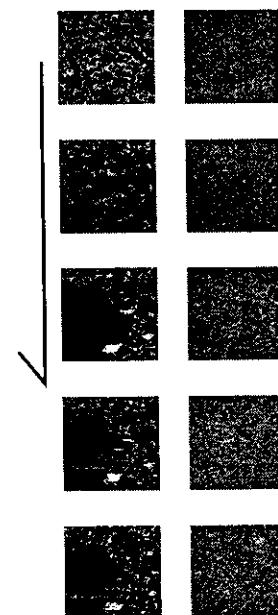
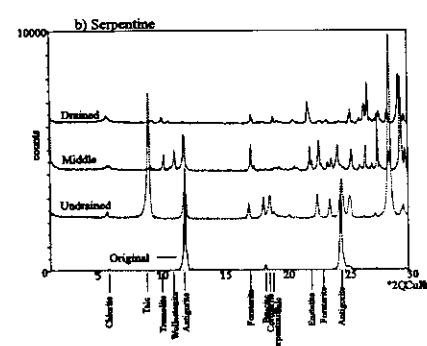
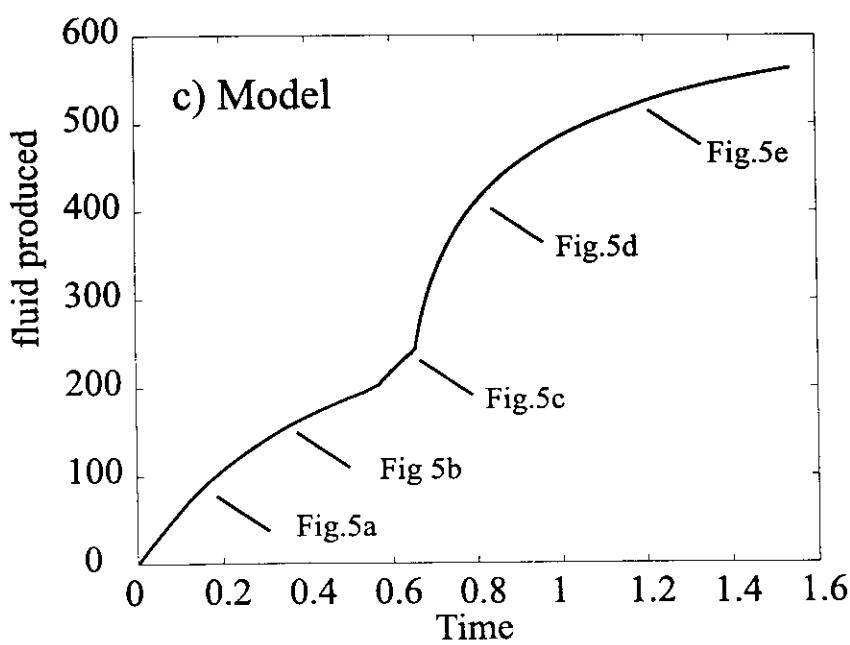
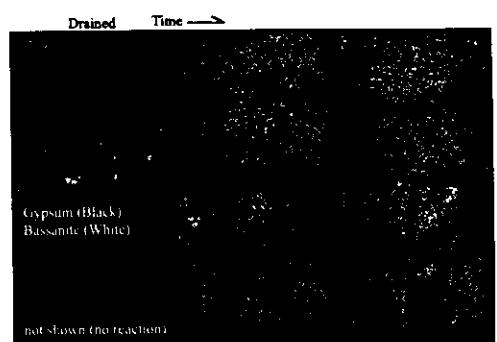
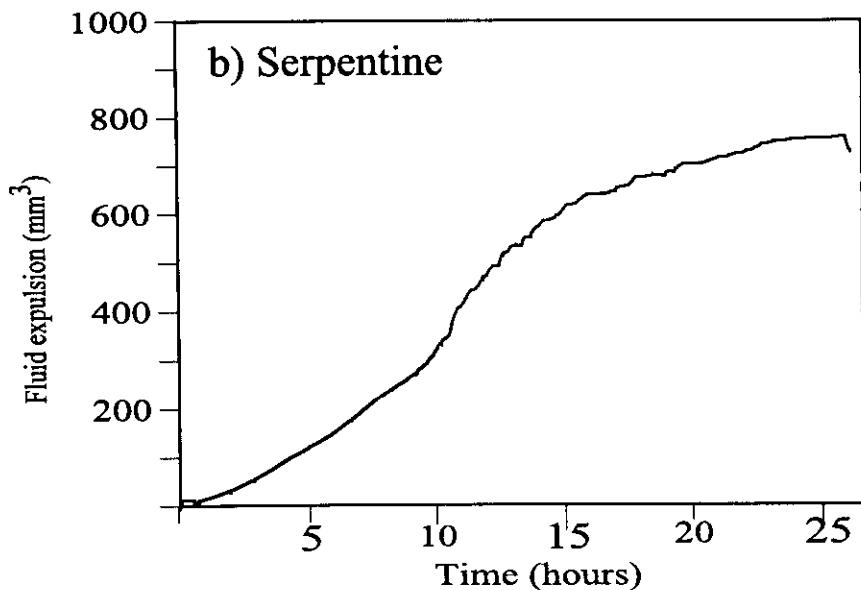
Model Results of Drained Boundary Connectivity and Fluid Production



Comparison of Experiments and Model



$\sim 1380 \text{ mm}^3$ total fluid



Conclusions

Simple model for hydrofracture-controlled systems shows:

- Self-organizing behavior.

- Produces power -law statistics of cluster connectivity at the critical state.

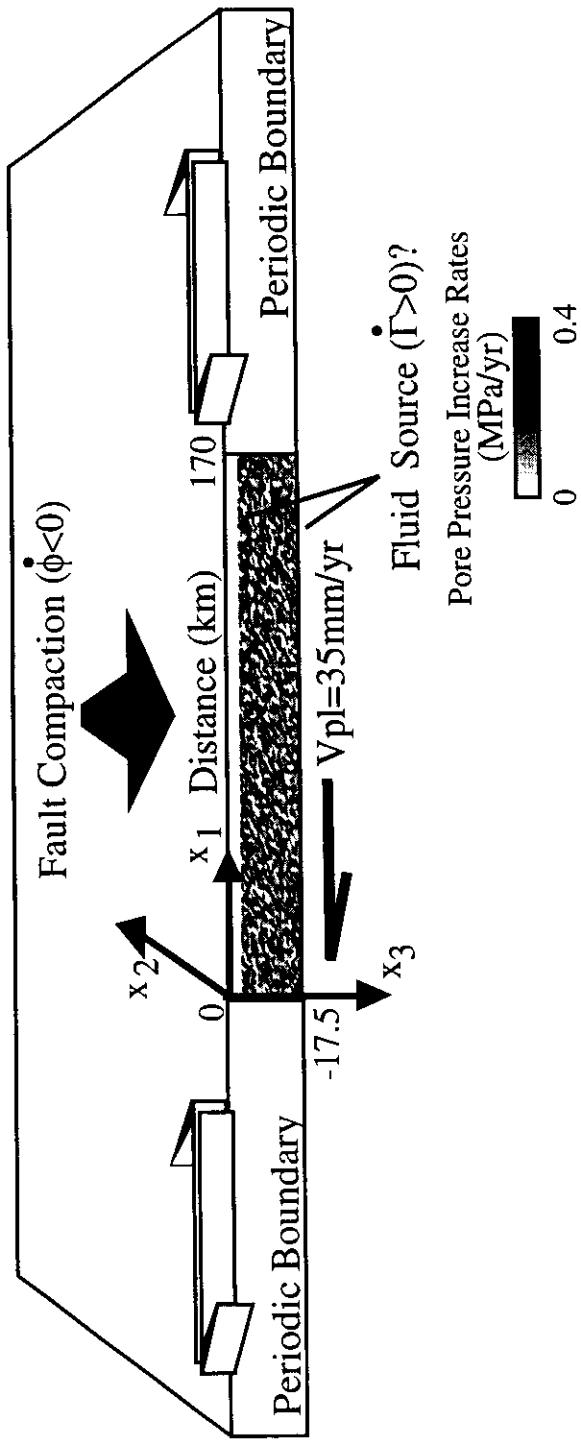
- Large (sudden) fluctuations in pore pressure from connectivity with lower pressure regions.

- Complex distribution of reaction products when coupled with kinetics of dehydration.

- Good agreement with fluid expulsion curves of dehydration experiments in serpentine and gypsum.

- Model applies to any kinetic system with positive Clapyron slopes.

3-D Fluid controlled earthquake model



Milner et al (2009)

Mohr-space dynamical system

