



INTERNATIONAL ATOMIC ENERGY AGENCY
UNITED NATIONS EDUCATIONAL, SCIENTIFIC AND CULTURAL ORGANIZATION



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WORKSHOP
GLOBAL GEOPHYSICAL INFORMATICS WITH APPLICATIONS TO
RESEARCH IN EARTHQUAKE PREDICTIONS AND REDUCTION OF
SEISMIC RISK

(15 November - 16 December 1983)

MODELLING SEISMICITY

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Why Model?

Some of
Earthquake sequences appear random, but the randomness may result from "simple" nonlinear mechanisms rather than true complexity.

Randomness

occurs to the extent that something cannot be predicted

is a matter of degree. Some systems are more predictable than others

often arises out of a sensitive dependence on initial conditions because the system amplifies noise.

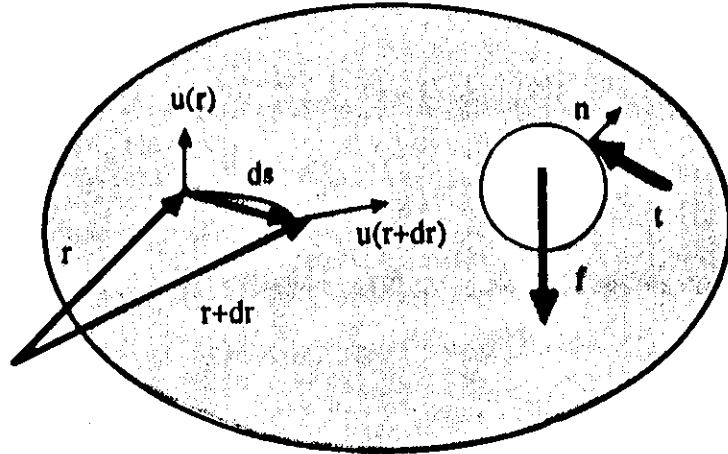
non linear

may be associated with attractors of varying dimensions

The theory of random processes is an empirical technique for coping with inadequate information and makes no statements about the causes of randomness.

The only truly fundamental source of randomness is the uncertainty principle of quantum physics.

Basic ideas of Continuum Mechanics



$$\int_S \mathbf{t} ds + \int_V \mathbf{f} dv = \int_V \rho \frac{d^2 \mathbf{u}}{dt^2} dv$$

$$\begin{bmatrix} t_x \\ t_y \\ t_z \end{bmatrix} = \begin{bmatrix} \sigma_{xx} & \sigma_{xy} & \sigma_{xz} \\ \sigma_{yx} & \sigma_{yy} & \sigma_{yz} \\ \sigma_{zx} & \sigma_{zy} & \sigma_{zz} \end{bmatrix} \begin{bmatrix} n_x \\ n_y \\ n_z \end{bmatrix}$$

Block Mechanics?

$$\begin{bmatrix} \frac{\partial u_x}{\partial x} dx + \frac{\partial u_x}{\partial y} dy + \frac{\partial u_x}{\partial z} dz \\ \frac{\partial u_y}{\partial x} dx + \frac{\partial u_y}{\partial y} dy + \frac{\partial u_y}{\partial z} dz \\ \frac{\partial u_z}{\partial x} dx + \frac{\partial u_z}{\partial y} dy + \frac{\partial u_z}{\partial z} dz \end{bmatrix}$$

$$\boldsymbol{\varepsilon} = \begin{bmatrix} \frac{\partial u_x}{\partial x} & \frac{1}{2} \left(\frac{\partial u_x}{\partial y} + \frac{\partial u_y}{\partial x} \right) & \frac{1}{2} \left(\frac{\partial u_x}{\partial z} + \frac{\partial u_z}{\partial x} \right) \\ \frac{1}{2} \left(\frac{\partial u_y}{\partial x} + \frac{\partial u_x}{\partial y} \right) & \frac{\partial u_y}{\partial y} & \frac{1}{2} \left(\frac{\partial u_y}{\partial z} + \frac{\partial u_z}{\partial y} \right) \\ \frac{1}{2} \left(\frac{\partial u_z}{\partial x} + \frac{\partial u_x}{\partial z} \right) & \frac{1}{2} \left(\frac{\partial u_z}{\partial y} + \frac{\partial u_y}{\partial z} \right) & \frac{\partial u_z}{\partial z} \end{bmatrix}$$

Finite Strain?

Constitutive Equations

1. Linear Elastic theories

$$\sigma_{xx} = \lambda \left(\frac{\partial u_x}{\partial x} + \frac{\partial u_y}{\partial y} + \frac{\partial u_z}{\partial z} \right) + 2\mu \frac{\partial u_x}{\partial x}$$

$$\sigma_{xy} = \mu \left(\frac{\partial u_x}{\partial y} + \frac{\partial u_y}{\partial x} \right)$$

$$\nabla \cdot \sigma = 0 \quad \longleftrightarrow \quad \vec{F} = m\vec{a}$$

2. Biot theory

$$\sigma_{ij} - \alpha \delta_{ij} p = \lambda \delta_{ij} \nabla \cdot \mathbf{u} + \mu \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right)$$

$$\nabla \cdot \sigma = 0$$

$$p = a_1 \theta + a_2 \nabla \cdot \mathbf{u} \quad \mathbf{v} = K \nabla p$$

$$\frac{\partial \theta}{\partial t} = -\nabla \cdot \mathbf{v}$$

Modify the principal stresses by a fraction of the fluid pressure. This is also a linear theory, originally

developed for soils. Notice we now have diffusion.

3. Temperature dependence See Aki and Richards page 20 ff

Rate of doing Mechanical Work + Rate of Heating =
= Rate of increase of (kinetic + internal)
energy

Rate of doing Mechanical Work =

$$= \int_{\mathbb{V}} \mathbf{f} \cdot \mathbf{v} d\mathbb{V} + \int_{\mathbb{S}} \mathbf{t} \cdot \mathbf{v} d\mathbb{S}$$

$$= \int_{\mathbb{V}} \mathbf{f} \cdot \mathbf{v} d\mathbb{V} + \int_{\mathbb{S}} \mathbf{v} \cdot \sigma \cdot \mathbf{n} d\mathbb{S}$$

$$= \int_{\mathbb{V}} \left(\rho \mathbf{v} \cdot \frac{\partial \mathbf{v}}{\partial t} + \nabla \mathbf{v} : \sigma \right) d\mathbb{V}$$

$$= \frac{d}{dt} \int_{\mathbb{V}} \left(\rho \frac{\mathbf{v} \cdot \mathbf{v}}{2} + \sigma : \dot{\epsilon} \right) d\mathbb{V}$$

= Rate of increase of kinetic energy +

?

$$\text{Rate of Heating} = - \int_S \mathbf{h} \cdot \mathbf{n} dS = \frac{d}{dt} \int_V Q dV$$

Rate of increase of internal energy per unit volume =

Rate of doing Mechanical Work + Rate of Heating -

Rate of increase of kinetic energy

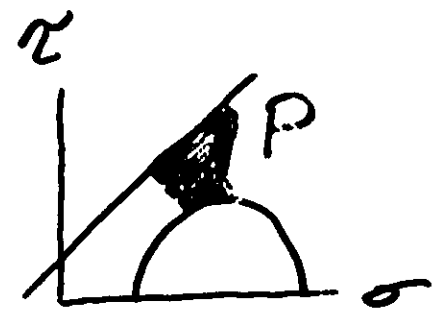
$$\frac{dU}{dt} = - \nabla \cdot \mathbf{h} + \sigma : \dot{\epsilon} = \frac{dQ}{dt} + \sigma : \dot{\epsilon} \text{ or}$$

$$dU = dQ + \sigma : d\epsilon = T dS + \sigma : d\epsilon$$

Idea $\oint du = 0$ for Equilibrium Thermodynamics

Is the occurrence of an earthquake to be described by equilibrium thermodynamics?

Probability of "Failure."



What is the form of P?

P_0

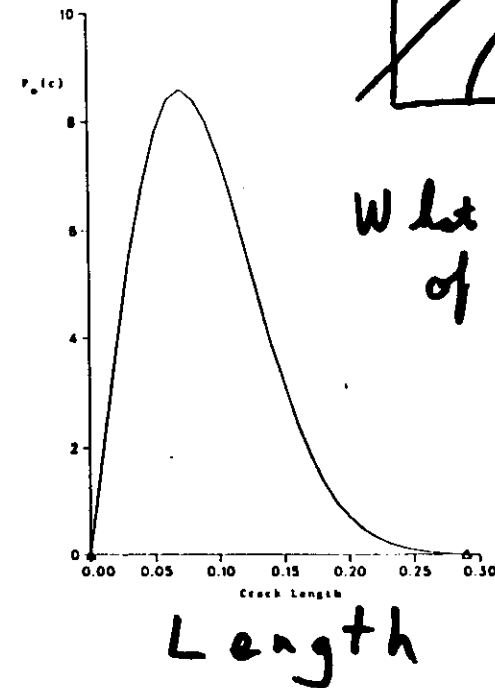


Figure 11... shows the assumed distribution of cracks in a body. $P_0(c)$ is calculated for $\sigma=0.1$.

M.Sc. Thesis

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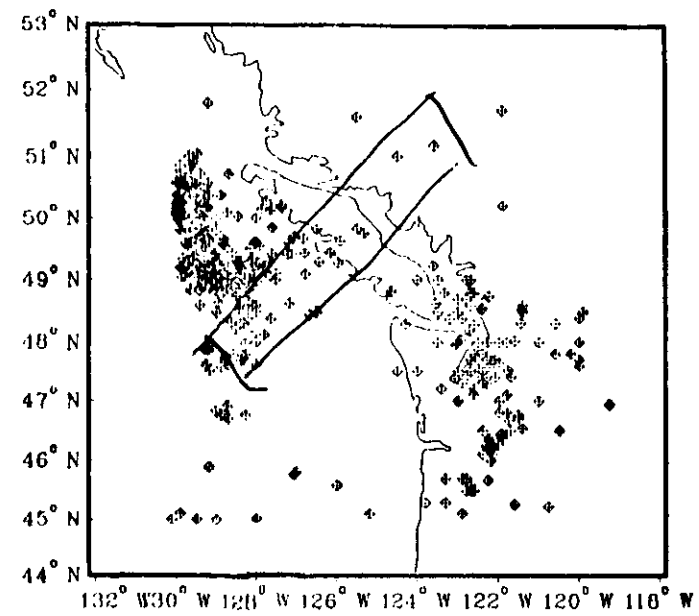
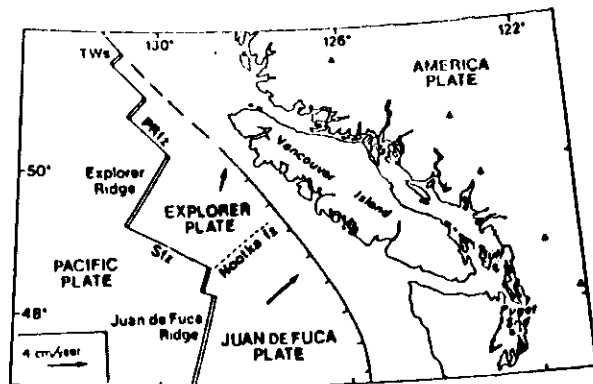


Figure 3

Earthquake Locations

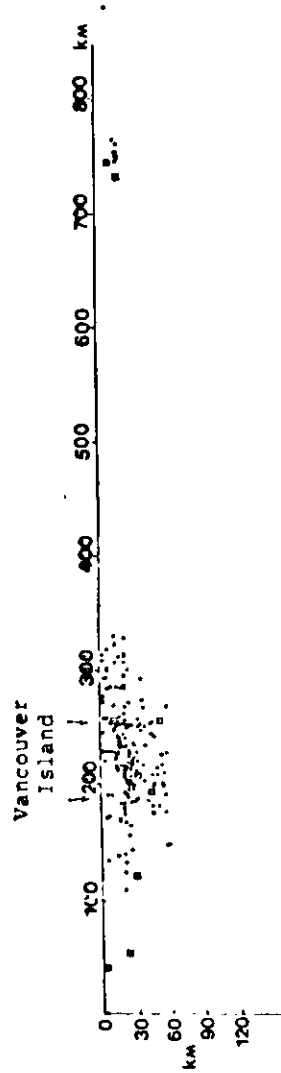
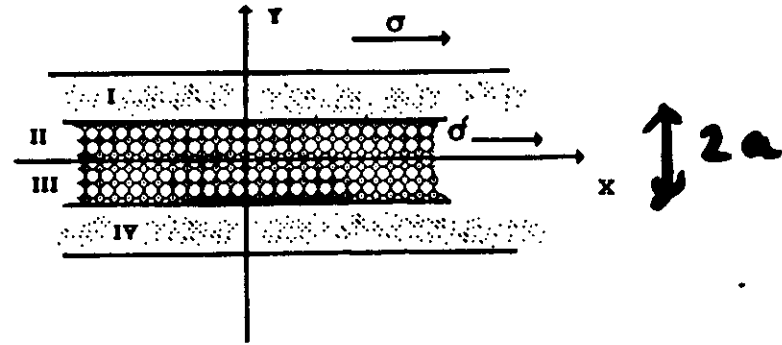
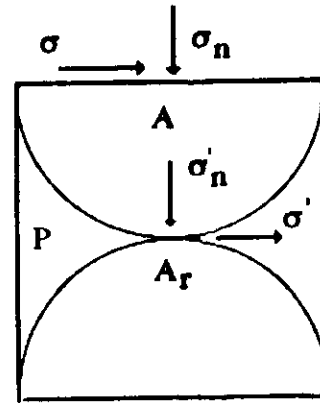


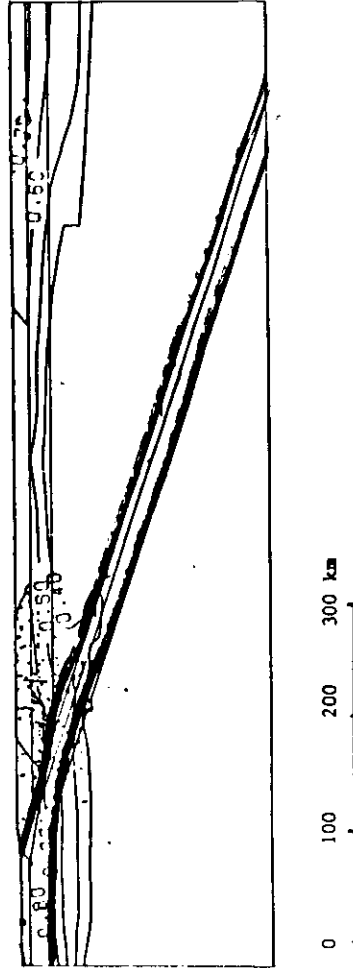
Figure 11.1 shows the distribution of the foci of the earthquakes within the rectangle in Figure 11.1.



If a fault zone is a mixture of "blocks" and fluids, what is the stress on the blocks?



σ is shear stress applied to total area A , σ' is the shear stress applied to the real area of contact A_r , P is pore pressure, σ_n is normal stress applied to A , σ'_n is normal stress applied to A_r



Push

Figure 15 shows the distribution of instability for the model with ridge push force exerted on oceanic plate.

$$\dot{\epsilon} = D_0 f(\tau) e^{\frac{AT_m}{T}}$$

$$\dot{\epsilon} = c_0^{-1} e^{\frac{-E}{kT}} \sinh \frac{v\tau}{kT}$$

$$\rho c_p \frac{\partial T}{\partial t} - \frac{\partial^2 T}{\partial x^2} = \tau \dot{\epsilon}$$

$$\rho c_p \frac{dT}{dt} = \tau \dot{\varepsilon} - 2 \frac{K}{a^2} (T - T_0)$$

$$\dot{\varepsilon} = \frac{1}{a} \frac{d\Delta}{dt}$$

$$\frac{\partial \Delta}{\partial t} = \frac{a}{c_0} e^{\frac{-E}{kT}} \sinh \frac{v \tau}{kT}$$

$$(\eta_{\text{eff}})_{\tau \rightarrow 0} = C^{-1} \frac{kT}{v} e^{\frac{E}{kT}}$$

$$v_0 \rho T e^{\frac{AT_m}{T}}$$

$$v_0 = B \left\{ \frac{1}{2} L_1 (1 + \tanh c' (T - T_m)) + L_2 \right\}^2$$

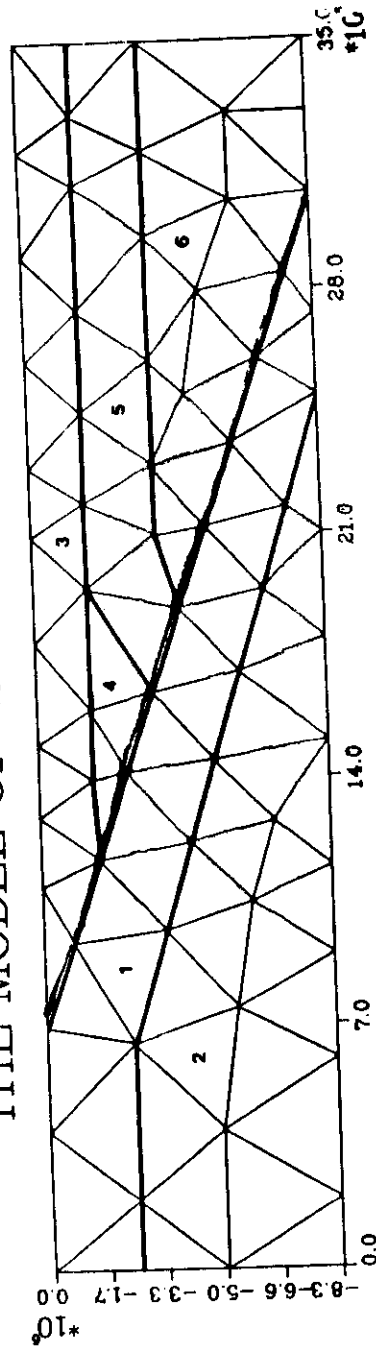
$$\frac{dL_1}{dt} = \alpha (L_1 - L_{\min}) \frac{\partial \Delta}{\partial t} - \frac{(L_1 - L_{\max})}{\tau_0}$$

$$\tau_0 = 10^6 e^{\frac{AT_m}{T}}$$

Displacement across a fault ~

Materials	E ($\times 10^{12} \text{ dyn/cm}^2$)	ν	ρ	Faults	E ($\times 10^{12} \text{ dyn/cm}^2$)	ν
1	1.2	0.25	3.3	1	2.0	0.4
2	1.04	0.3	2.295			
3	0.89	0.25	2.92			
4	1.163	0.3	3.3			
5	0.71	0.3	2.92			
6	1.25	0.3	3.285			

THE MODEL OF VANCOUVER ISLAND



$$\mathbf{Kd} = \mathbf{f} = \mathbf{A}_1 \mathbf{b} + \mathbf{A}_2 \delta + \mathbf{A}_3 \mathbf{g}$$

$$\boldsymbol{\tau} = \mathbf{M}_1 \mathbf{d} + \mathbf{M}_2 \Delta$$

$$\frac{d\boldsymbol{\tau}}{dt} = \mathbf{M}_1 \mathbf{K}^{-1} \left[\mathbf{A}_1 \frac{d\mathbf{b}}{dt} + \mathbf{A}_2 \frac{d\Delta}{dt} \right] + \mathbf{M}_2 \frac{d\Delta}{dt}$$

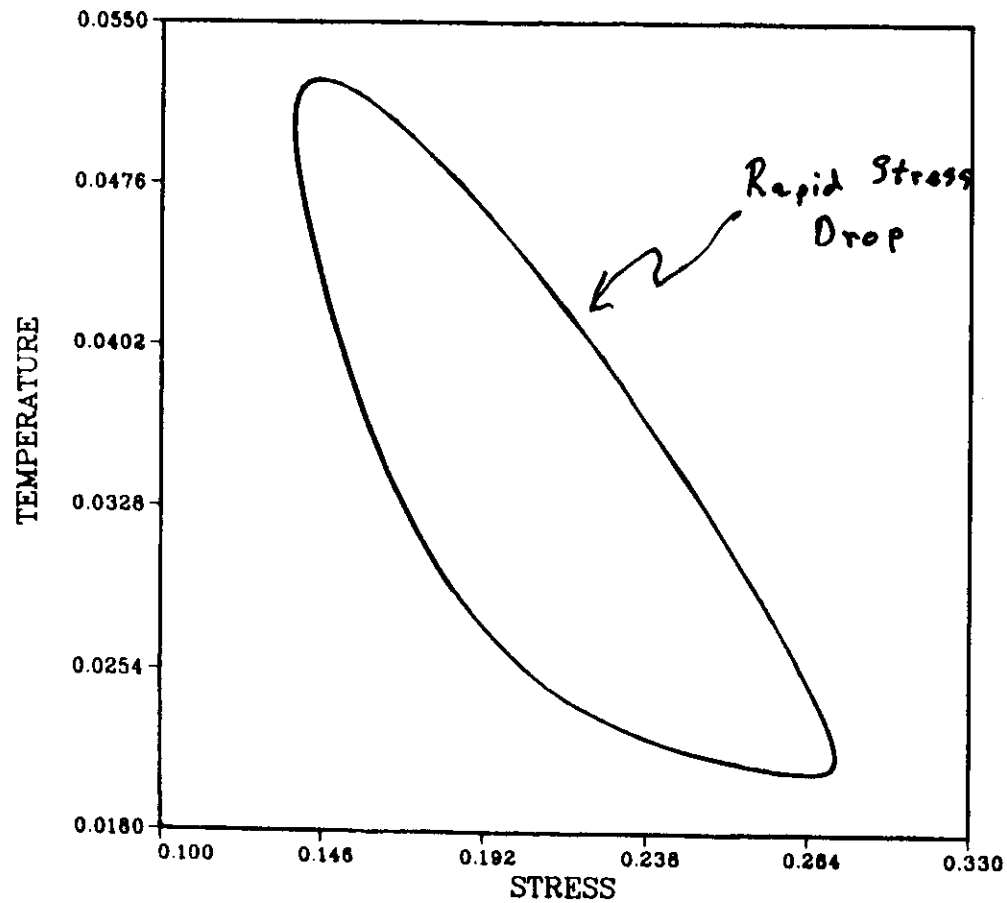
$$\mathbf{b}(t) = \mathbf{b}_0 + t \sum_{k=1}^K v_k \mathbf{c}_k = \mathbf{b}_0 + t \mathbf{C} \mathbf{v}$$

$$\frac{\partial \boldsymbol{\tau}}{\partial t} = [\mathbf{M}_1 \mathbf{K}^{-1} \mathbf{A}_1 \mathbf{C}] \mathbf{v} + [\mathbf{M}_1 \mathbf{K}^{-1} \mathbf{A}_2 + \mathbf{M}_2] \frac{d\Delta}{dt}$$

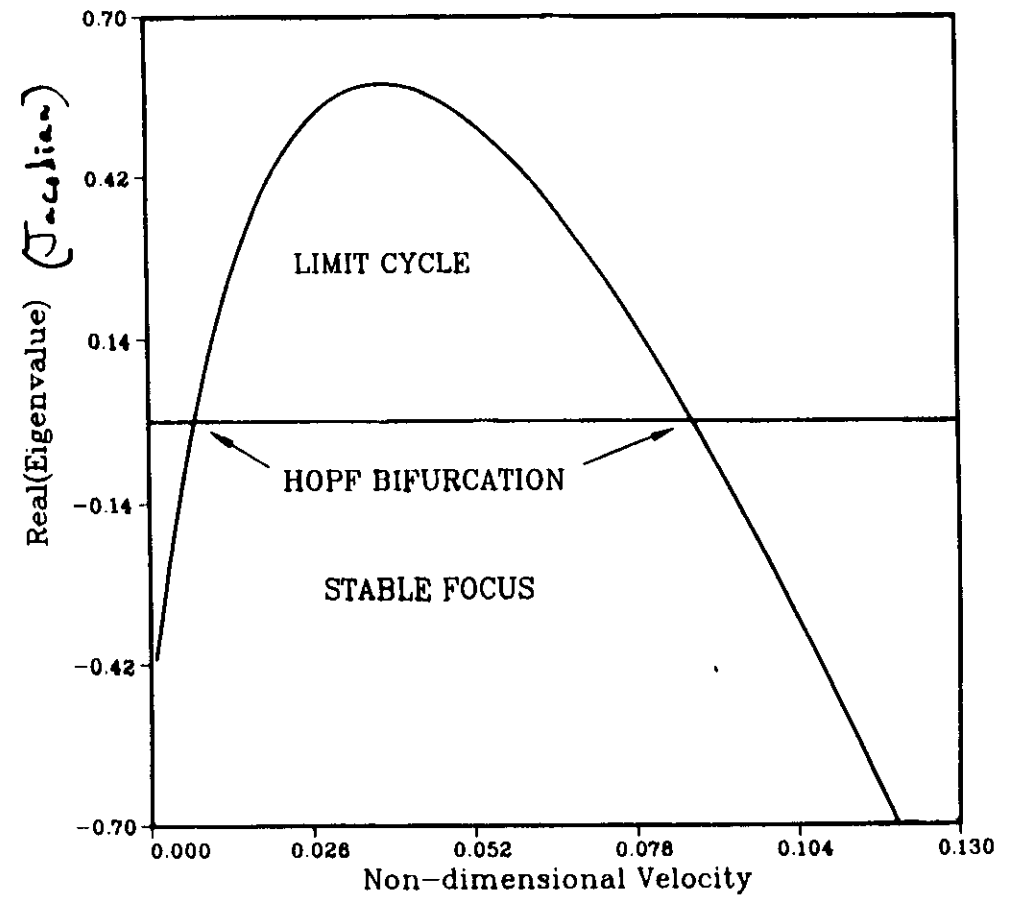
$$\mathbf{K}(\mathbf{K}^{-1} \mathbf{A}_1 \mathbf{C}) = \mathbf{A}_1 \mathbf{C} \quad \mathbf{K}(\mathbf{K}^{-1} \mathbf{A}_2) = \mathbf{A}_2$$

Solve a small number of FE problems

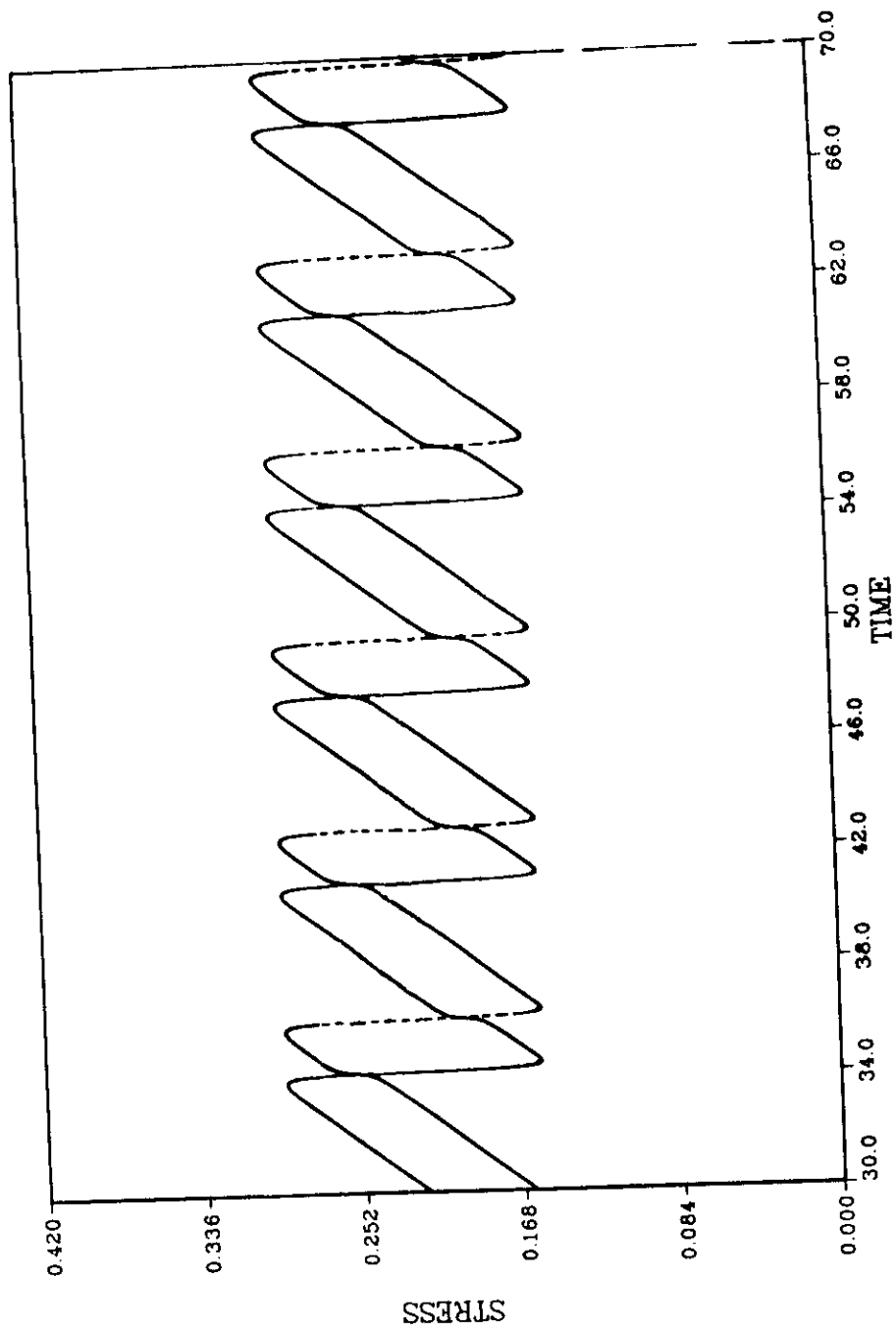
One Element System Phase Plot



One Element System Stability Analysis

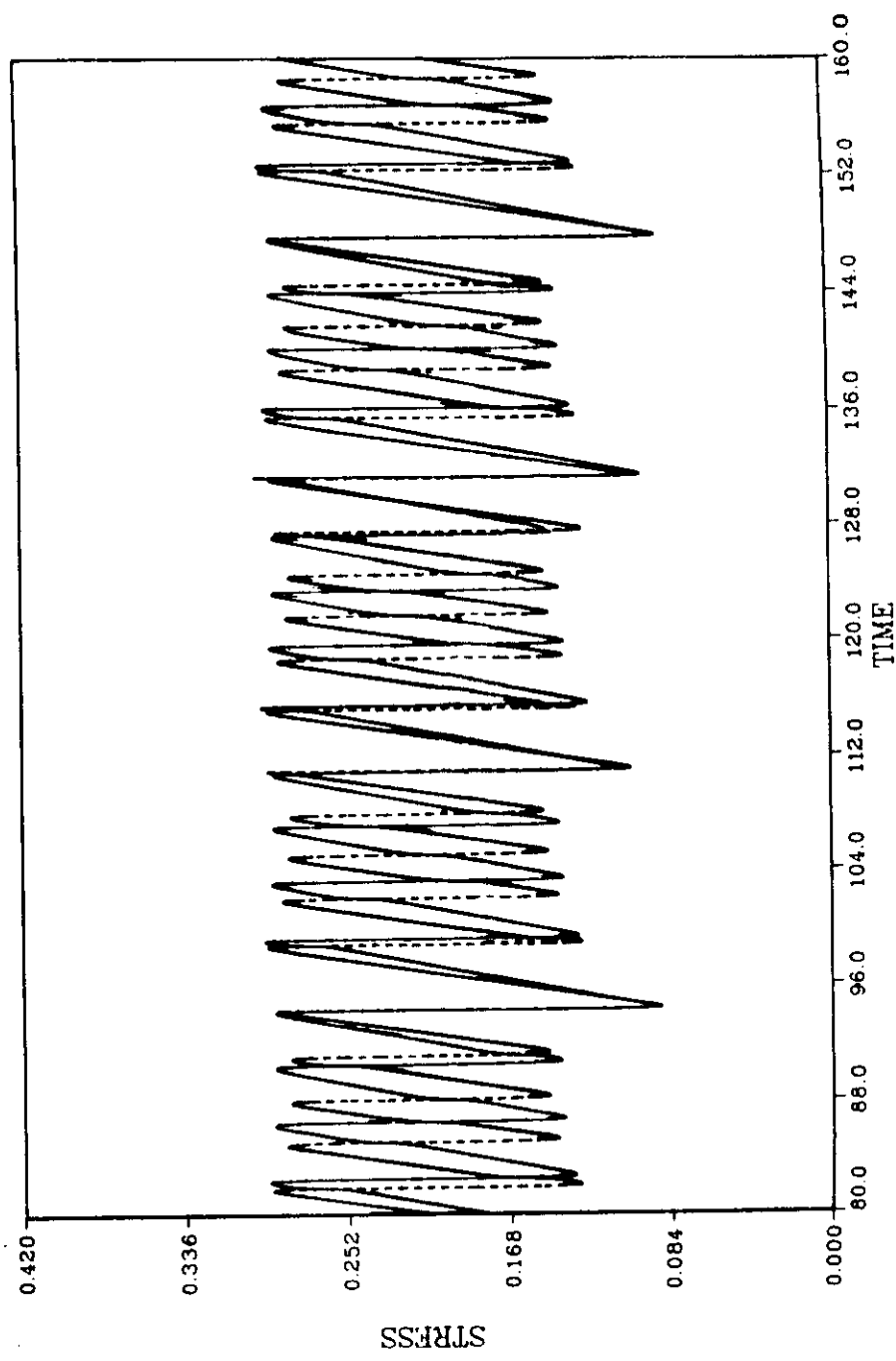


TWO ELEMENTS, EQUAL DRIVES

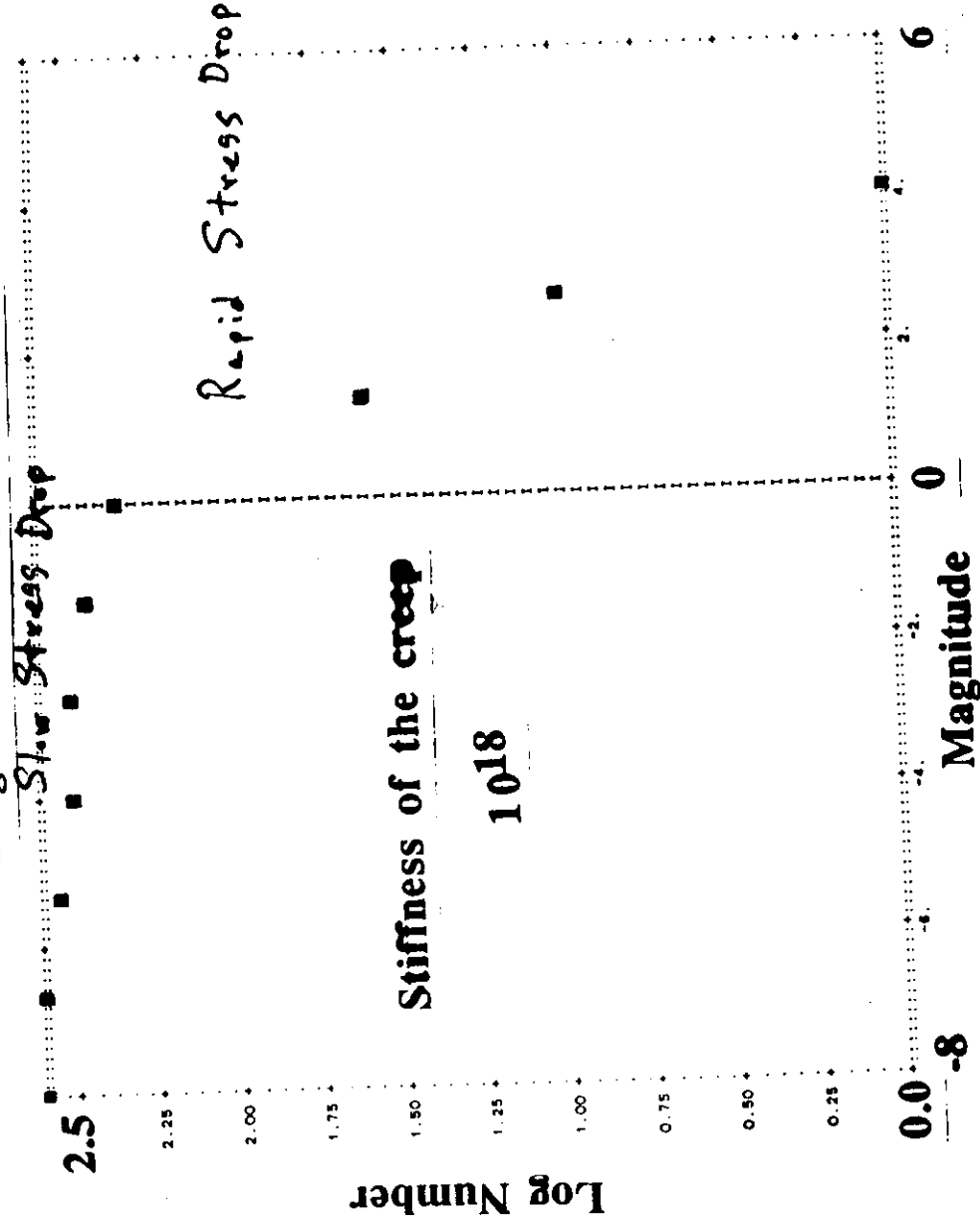


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TWO ELEMENTS, UNEQUAL DRIVES Effect of Breaking Symmetry

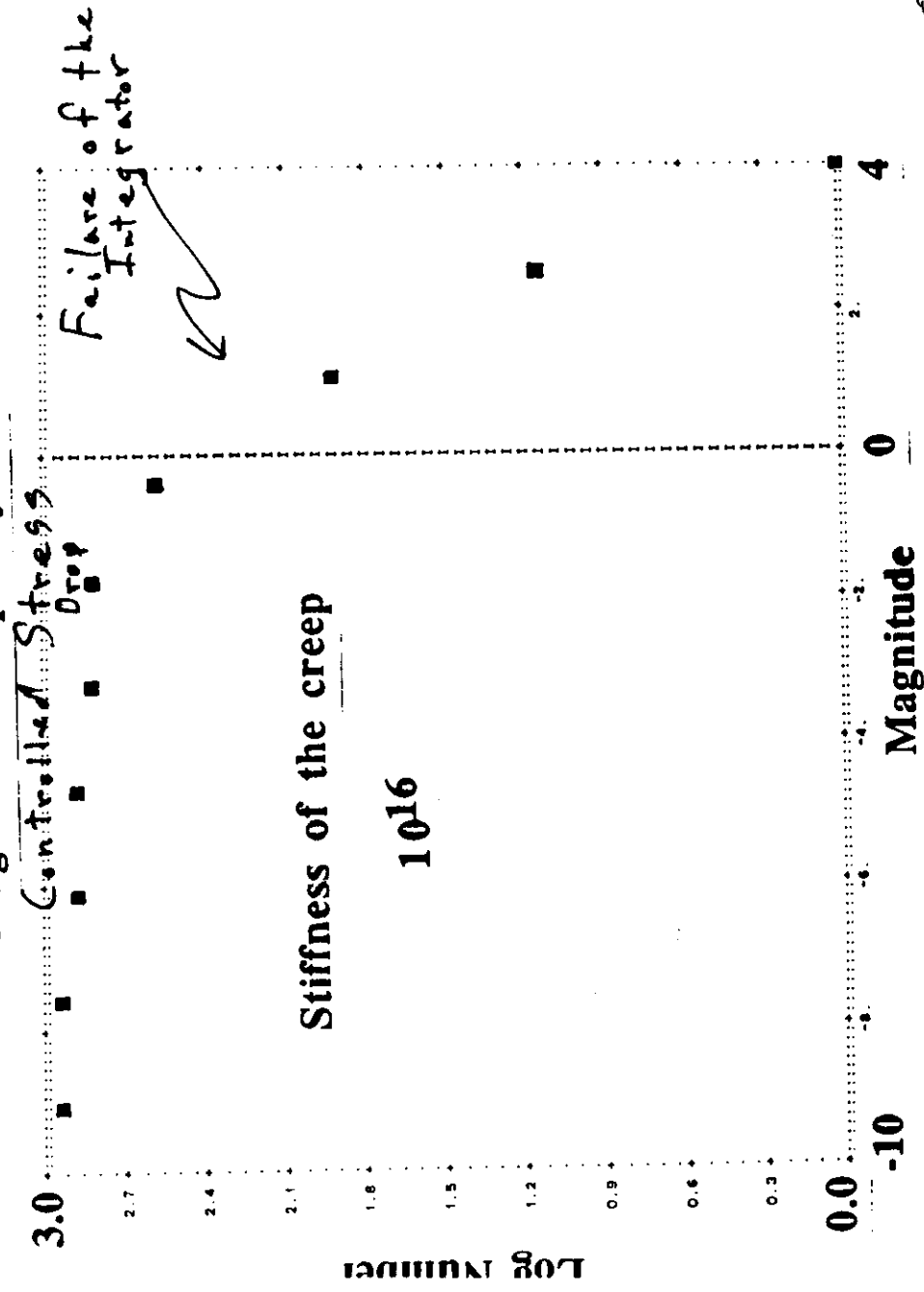


Magnitude Frequency Plot



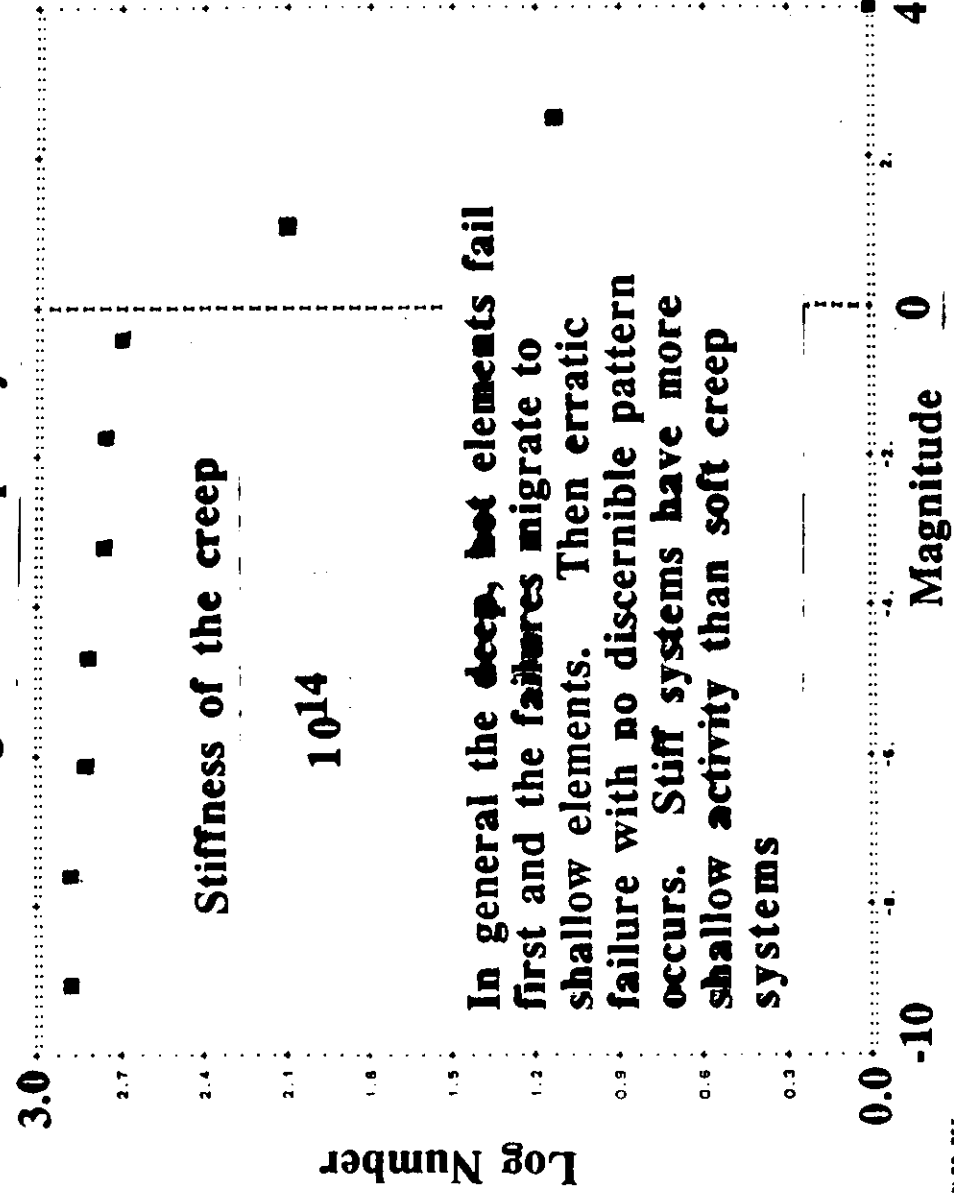
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Magnitude Frequency Plot



CG DFF

②



#LOG DFF

