



INTERNATIONAL ATOMIC ENERGY AGENCY
UNITED NATIONS EDUCATIONAL, SCIENTIFIC AND CULTURAL ORGANIZATION



INTERNATIONAL CENTRE FOR THEORETICAL PHYSICS
34100 TRIESTE (ITALY) • P.O.B. 506 - MIRAMARE - STRADA COSTIERA 11 • TELEPHONE: 2240-1
CABLE: CENTRATOM - TELEX 460862-1

H4.SMR/303 - 24

**WORKSHOP
GLOBAL GEOPHYSICAL INFORMATICS WITH APPLICATIONS TO
RESEARCH IN EARTHQUAKE PREDICTIONS AND REDUCTION OF
SEISMIC RISK**

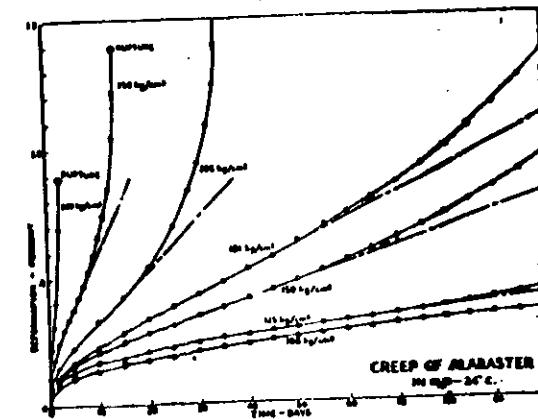
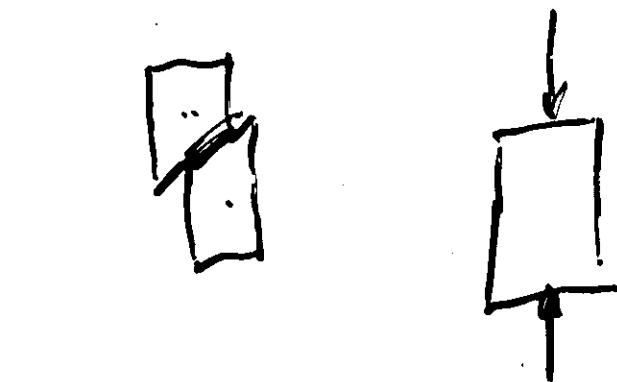
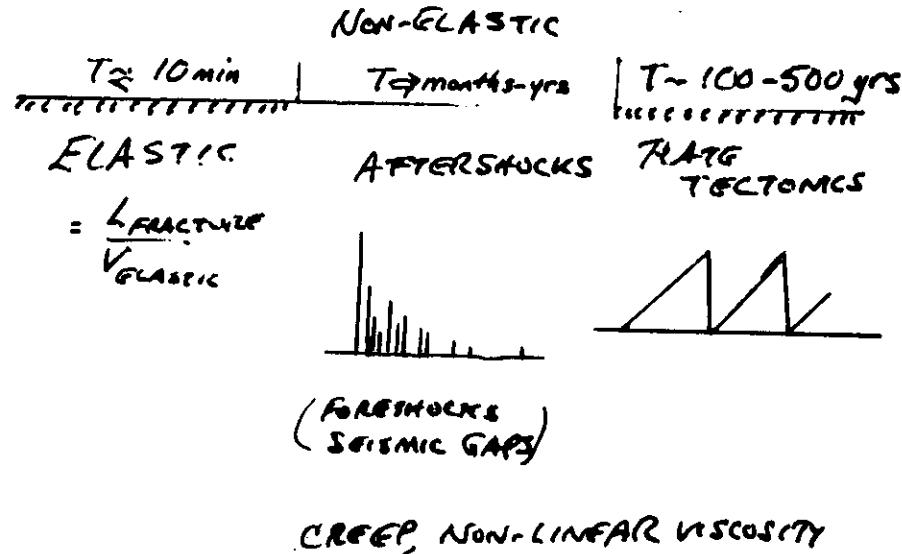
(15 November - 16 December 1988)

**LABORATORY CONSTRAINTS ON PHYSICAL MODELS
THE PHYSICS OF FRACTURE**

L. KNOPOFF

**University of California
Los Angeles
U.S.A.**

TIME SCALES



Chemical Kinetics
Activation Process $\rightarrow t_p = t_0 e^{-E/kT} \cdot E \approx 0 \text{ eV}$

Stress Corrosion
(role of H₂O)
a large number of references:

500

⑧

40

Glass : Preston et al, Nature 156, 1946
J. App. Phys. 17, 1946

Metals : Zimkov, Int Fract Mech, 1, 1965

Rust & D : Griggs, Bull. U.S.A., 51, 1940

Minerals, Ionic Solids: Zimkov, loc. cit.

Metals: Cherepanov, Mechanics of Brittle Fract. (McGraw-Hill) 1979.

TABLE I.—Properties of alabaster in wet creep tests at different stresses
(Gypsum in rock experiment subject to constant load, in the presence of its own saturated aqueous solution, at $T = 11^{\circ}\text{C}$, $v_1 = v_2 = 0$)

Compressive stress ($\sigma_1 - \sigma_2$, in kg/cm ²)	Steady-state velocity (millions per day) v_1	Deformation before fracture (per cent) δ	Duration of test (to fracture) t _f (days)	Equivalent viscosity (poises) $\times 10^4$
300	2000	.59	2.45	.42
250	440	1.49	13.88	1.60
208	218*	2.1*	45*	2.04
181	100	3.0	133	5.11
165	77.0	3.3	285	6.04
150	66.5	>1.05	>110	6.36
125	24.5	3.3	208	14.4
108	3.0*	>2.5*	>300*	41.0 (1)*
	7.6	>6.1	71.9 (2)	7.5

* Average of three duplicate tests.

Properties of same gypsum when dry:

Strength (short-time) = 550 kg/cm²

Deformation before fracture = 38% (primarily elastic)

Equivalent viscosity = $> 5 \times 10^4$ at 420 kg/cm²

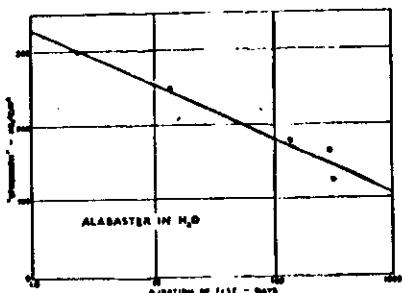
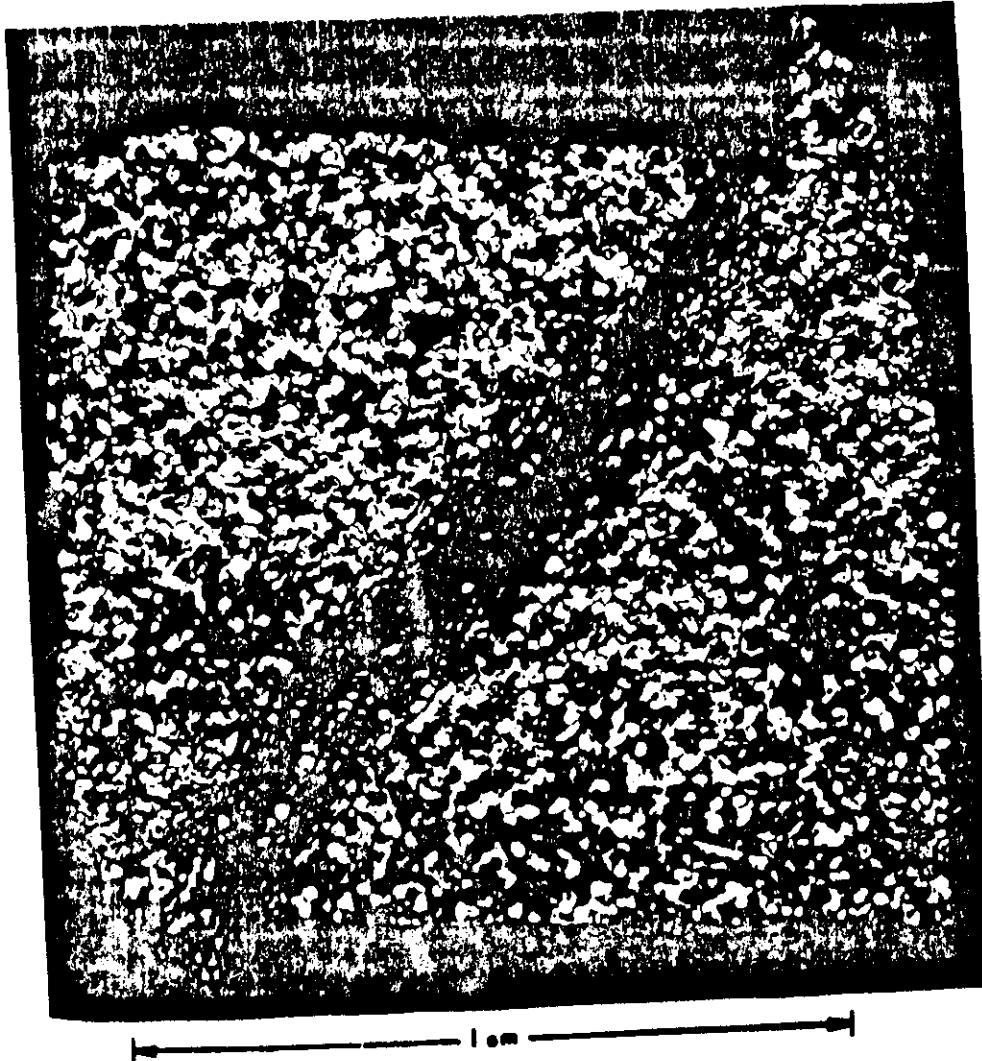
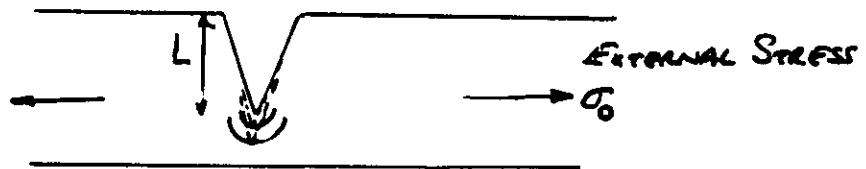


FIGURE 10.—Strength of alabaster as a function of duration of test

Data from experiments of Figure 8.



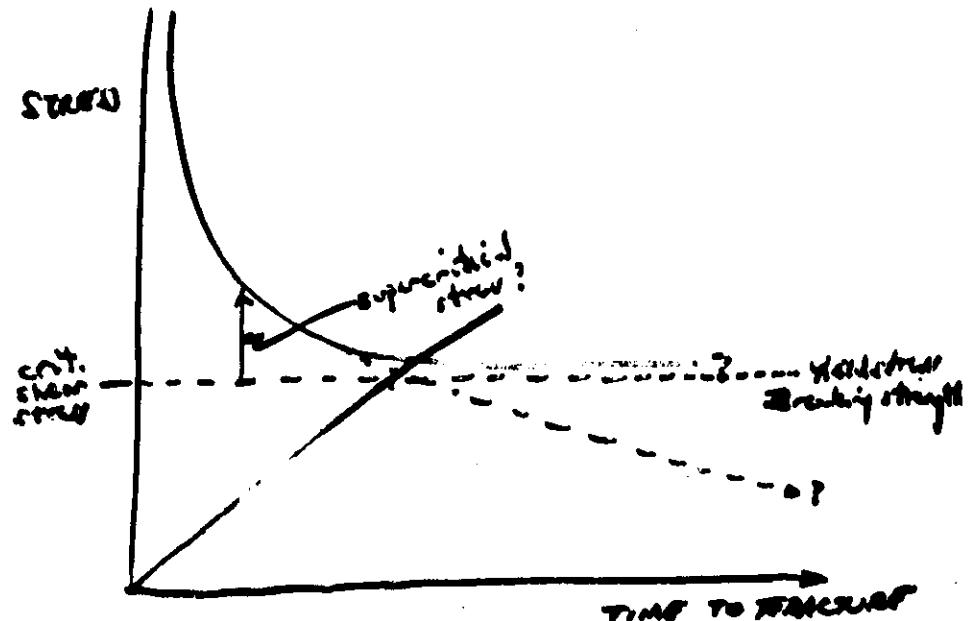
ST. PETER SAND 40-65 MESH
Compacted at 500°C., 5 kb confining pressure, 1 kb interstitial water pressure, then compressed 40 per cent.
Temperature and fluid pressures kept constant. Thick copper jacket did not rupture despite shear offset at ends.



$$\bullet V = \frac{dL}{dt} \sim e^{\sigma_0 \sqrt{L}/kT} \times \text{const}$$

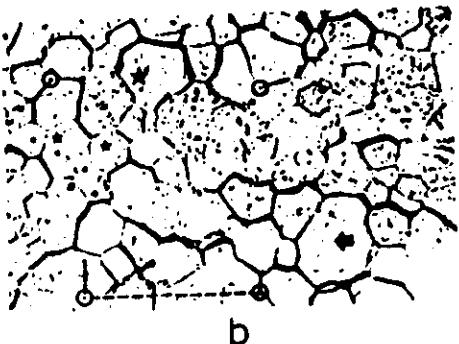
activation process!

- Process much accelerated in presence of water
IN ROCKS
SOLUTION & RICKEY STALLIZATION





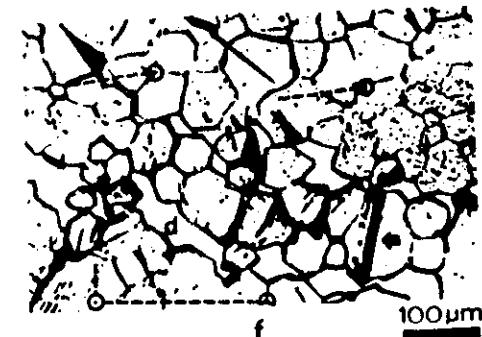
a



c



e

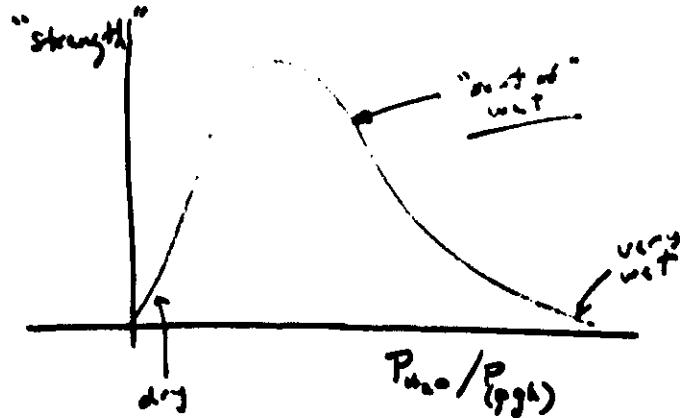


The Role of Water:

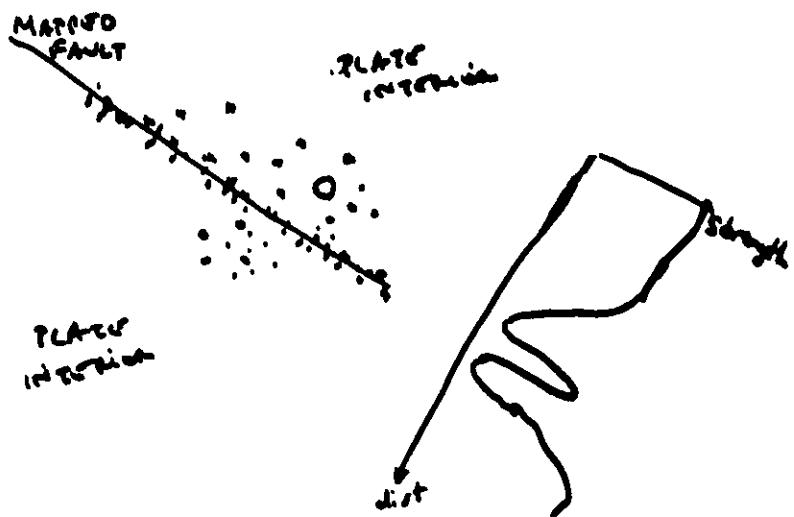
1. Why does wet sand have cohesive strength?
surface tension



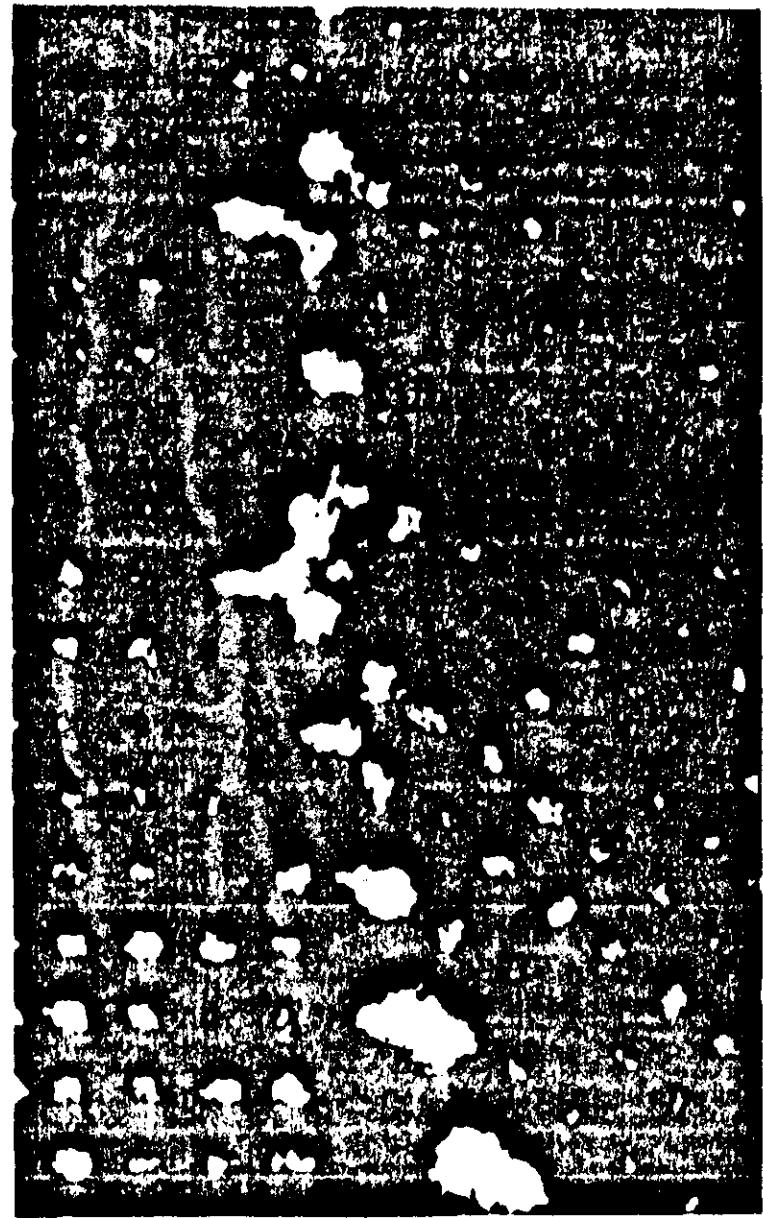
Figure 4. Cracking and twinning in sodium nitrate in ductile shear zone. Intergranular crack propagation and twin inception near small stars in b and d; twin widening at stubby arrow; microfracture dilation at d and elsewhere. Sample thickness is about 30 μ m.



1. experiments on
gouge
2. diffusion Mylonite

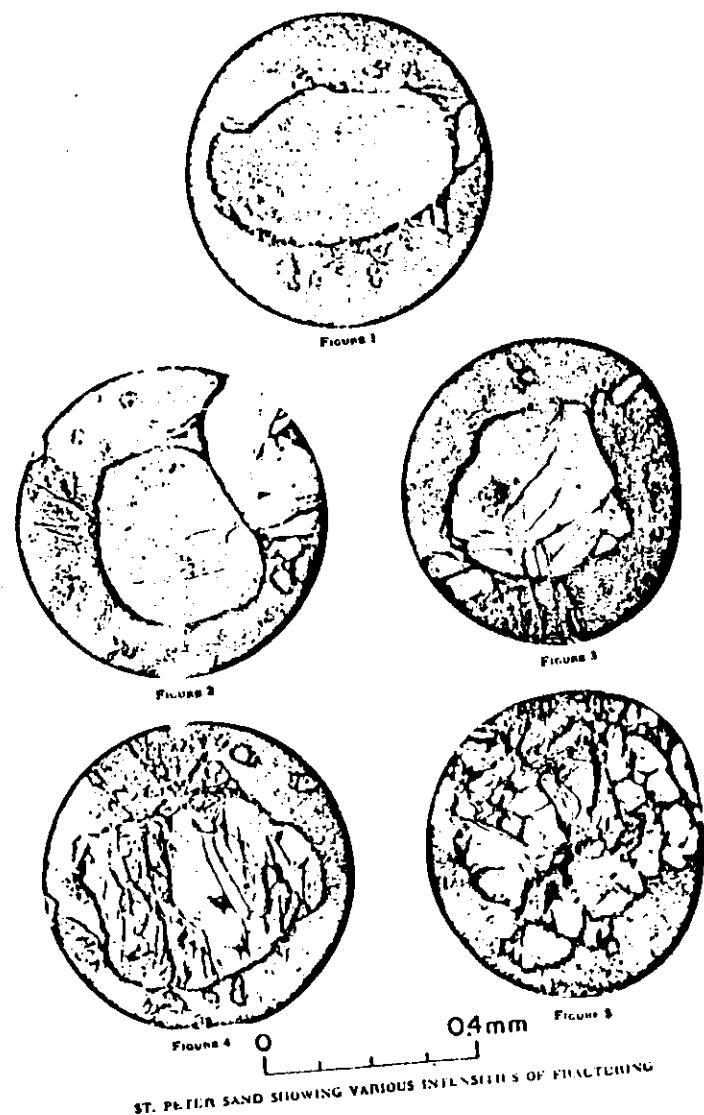


Faceted NIO boundary: The white areas represent the areas between atom columns and show the relatively wide spaces at the grain boundary. The segment of the 38° boundary on the left is asymmetric with the grain boundary parallel to a (100) plane in the lower half of the bicrystal. The segment on the right is approximately symmetric.

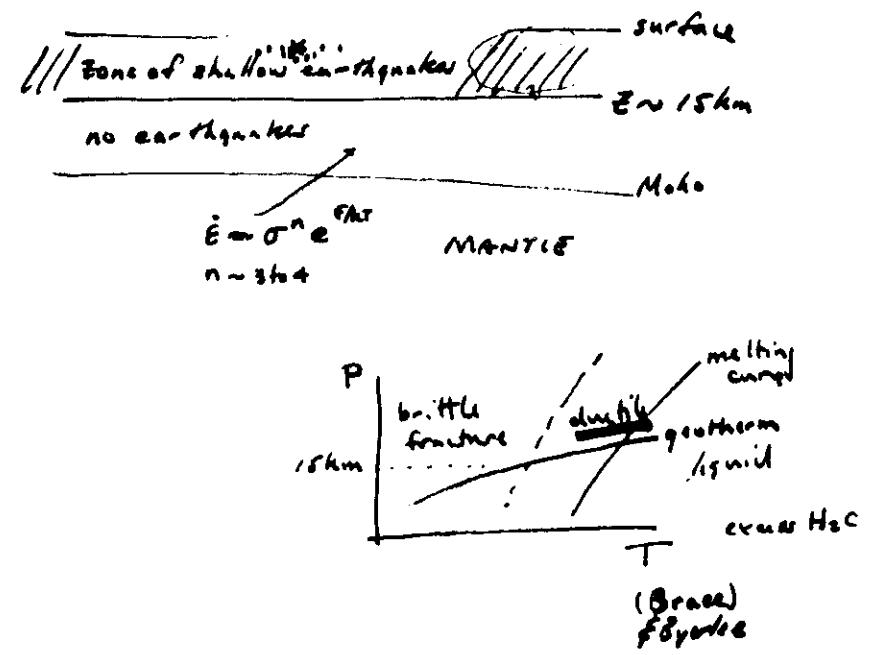


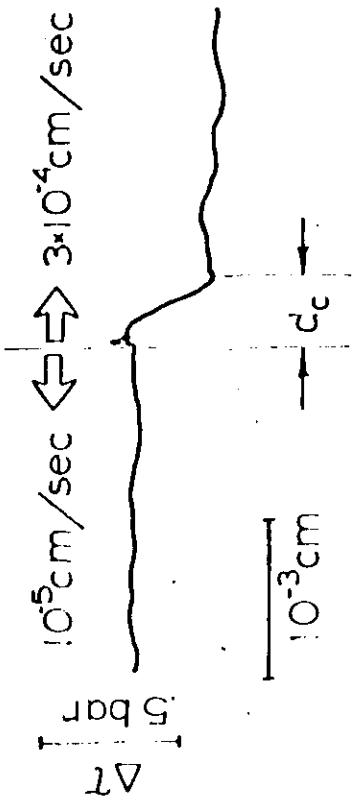
K. L. Merkle, Argonne National Laboratory

281



Is the viscosity of the earth everywhere and the same nature?

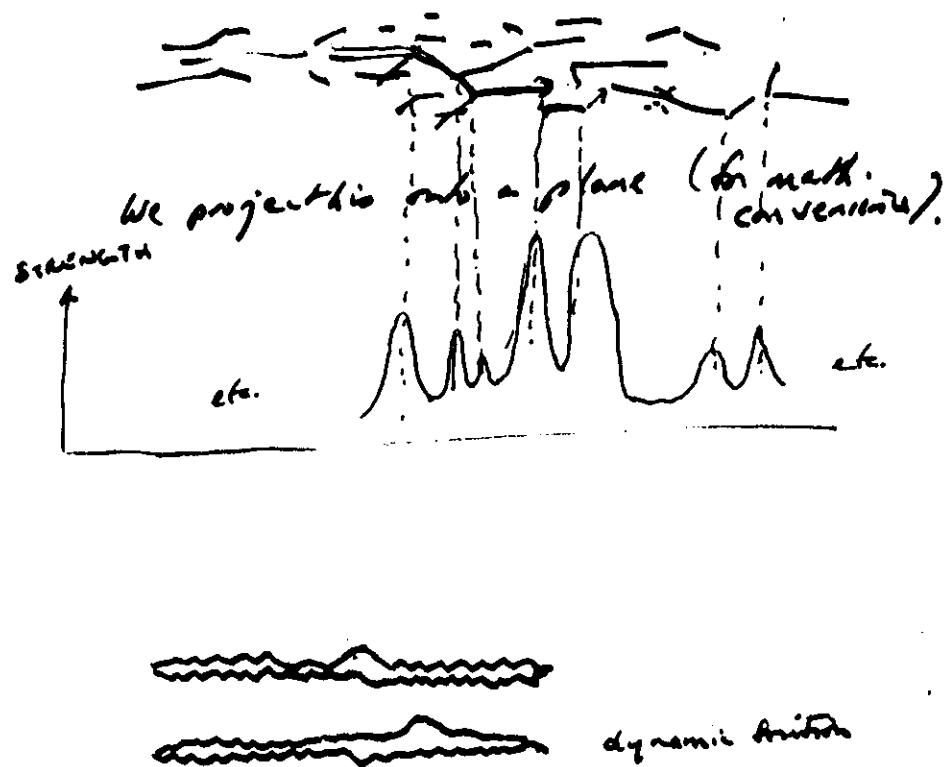




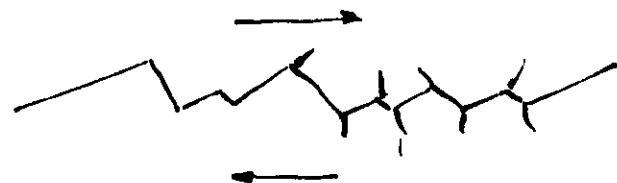
effect of kinetic factor
 $\propto \rho_0 \cdot \log(1+\alpha t)$
 Districh
 JGR (1979)

Figure 5

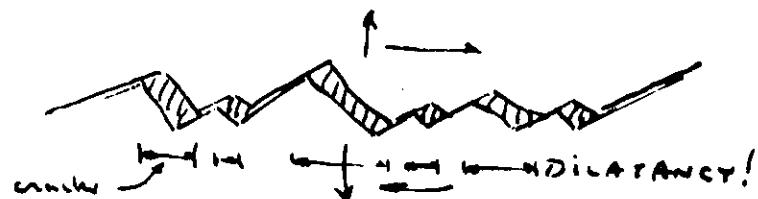
What is the yield strength of a barrier
 to and if the strength depends on
 geometry?



consider an irregular crack surface which we model as a sawtooth. We subject the elastic medium to a shear stress.



In order that lateral motion takes place, one crack surface must "ride" upward over the other. Thus transverse motion must take place

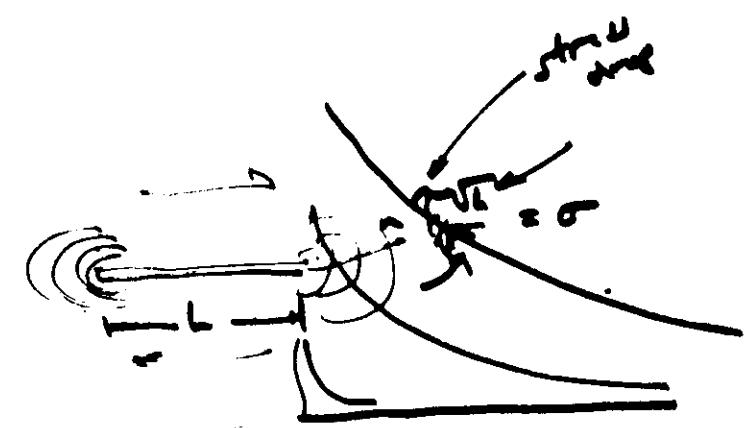
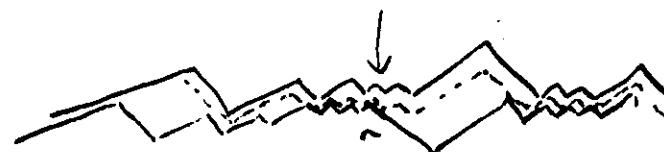


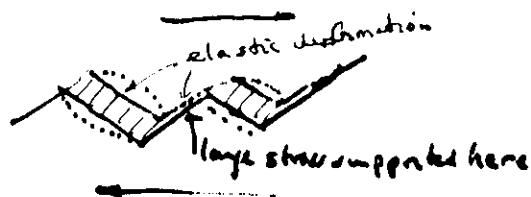
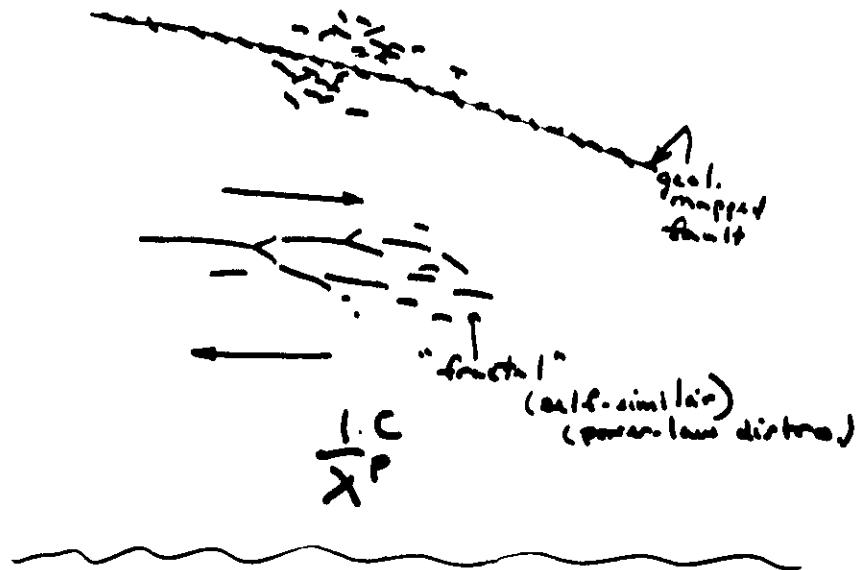
When sliding becomes too large, cracks begin to link up.



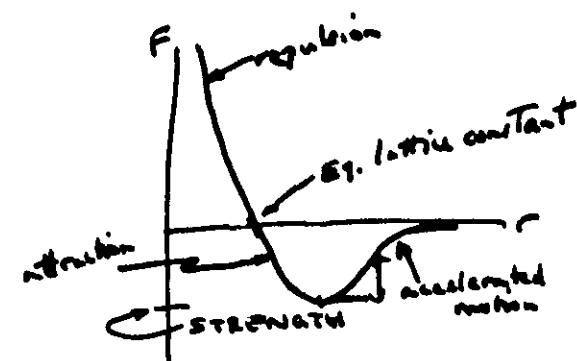
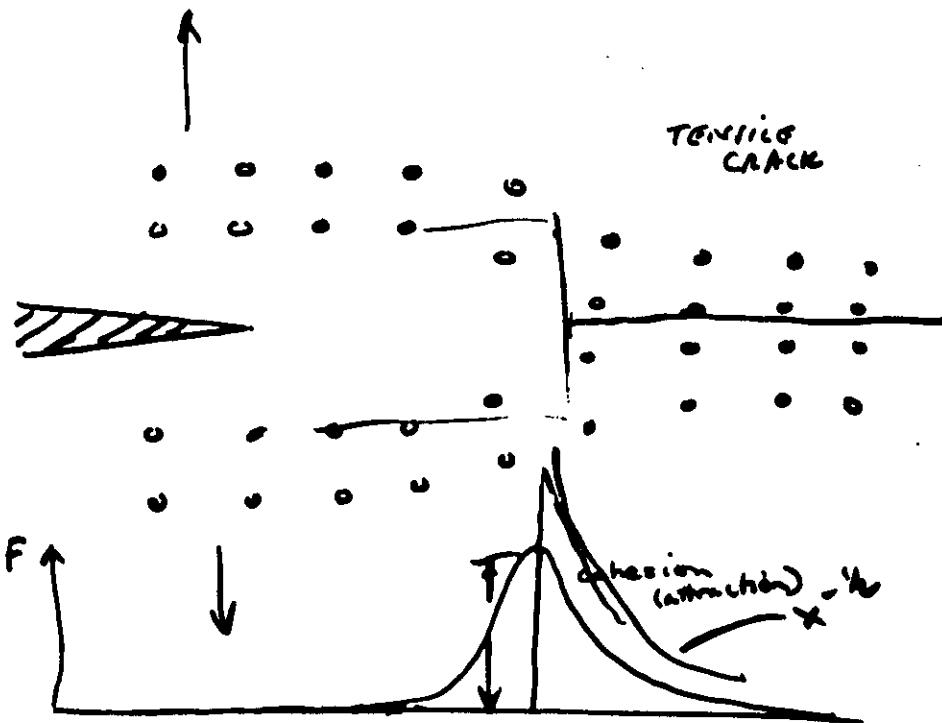
What keeps a crack "open"?

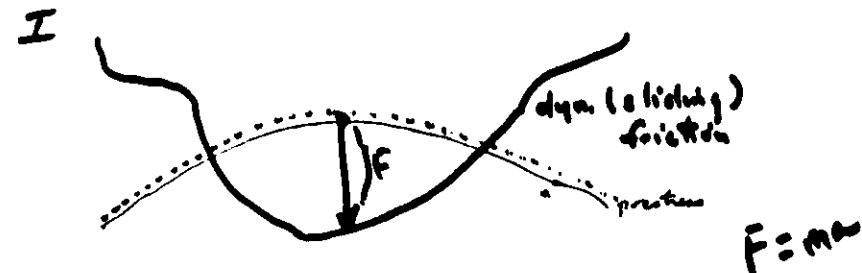
Heating!



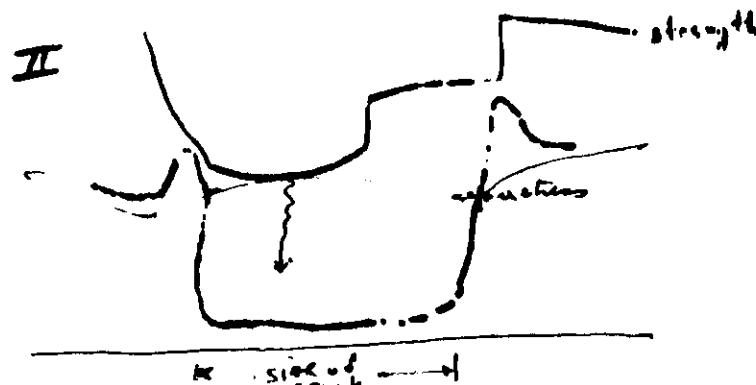


(role of water in
recrystallization)
what are angles of the
saw-teeth?
how high are the saw-teeth?
(wave length)

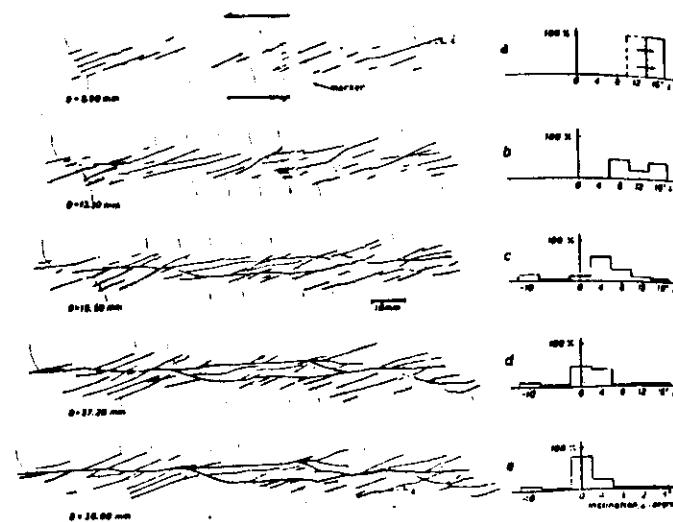
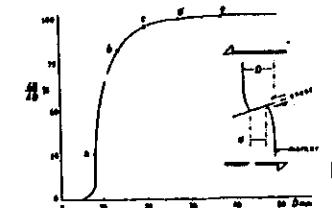
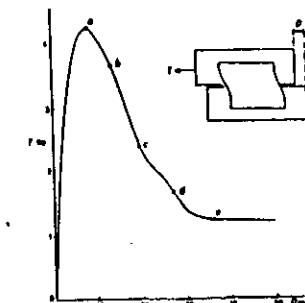
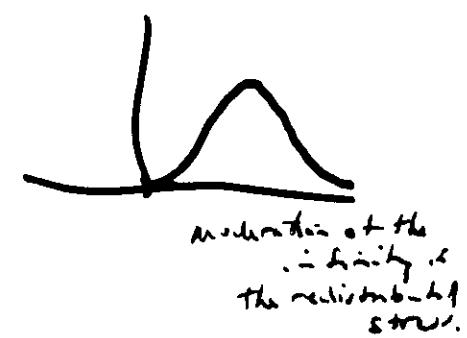




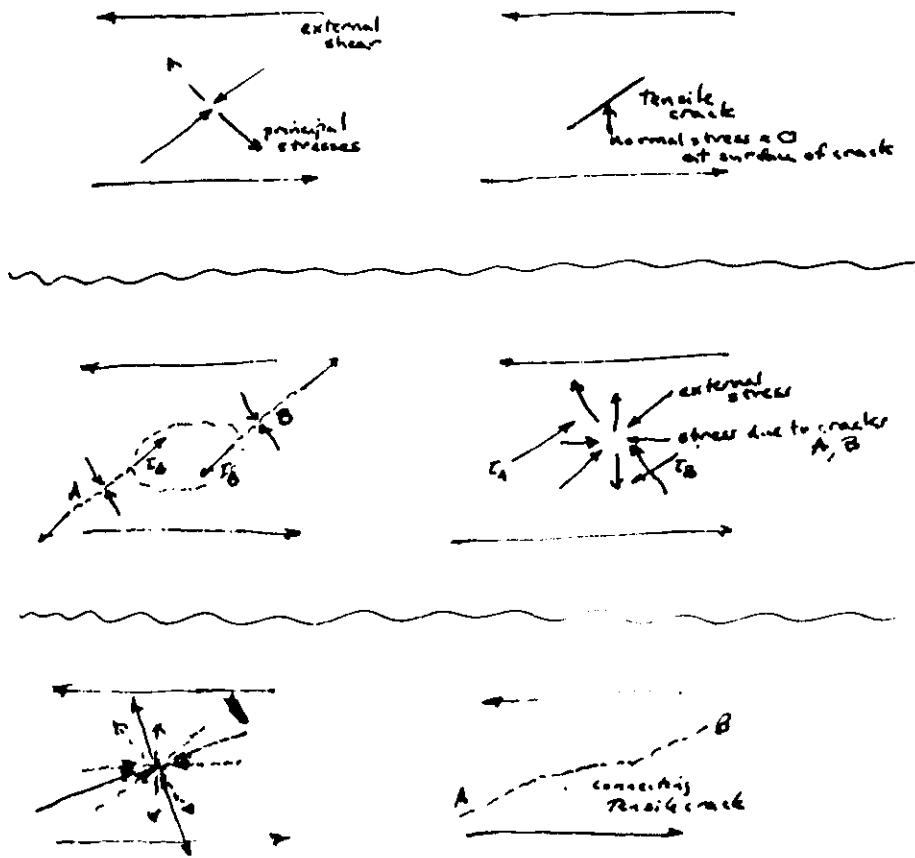
size of crack
"ever-increasing friction"
Burridge & Helfand
Geophys J.
R.A.S.
1972±1



"the unbreakable obstacle"

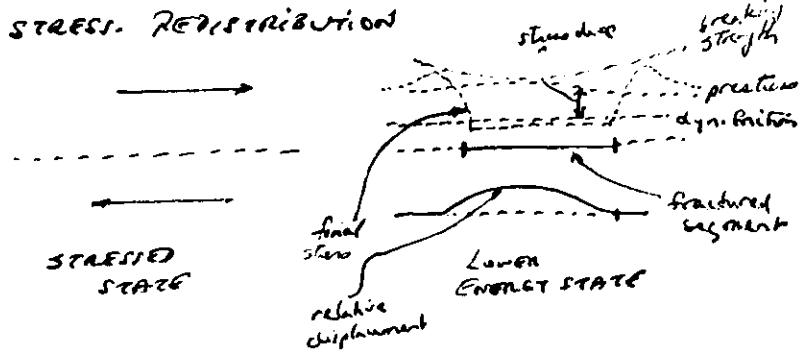


Tchalenko



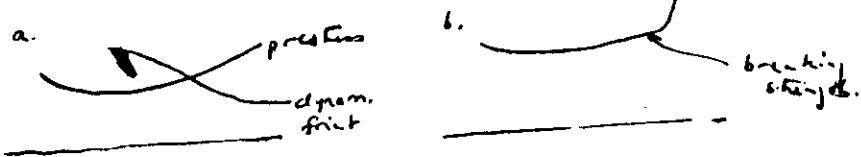
PHYSICS OF FRACTURE (simplified version)

1. A FRACTURE IN A PRESTRESSED SOLID LEADS TO A STRESS RE DISTRIBUTION



2. A fracture stops because

- a. It enters a region where the stress drop becomes negative
- b. It encounters a region of very large strength



I personally believe that model b) is the mechanism
(the barrier or viscosity model).

3. CREEP IS AN IMPORTANT ACCOMPLIMENT TO THE EARTHQUAKE EVENT

This process must involve some sort of (non-linear) rheology. The TIME INTEGRAL OCCURRED AFTERSHOCKS IS TOO SHORT TO BE ACCOUNTED FOR BY PLATE TECTONICS. PRE- AND POST-SHOCK CARGO IS OBSERVED IN THE FIELD.

An attack on the basic assumption
(that we know how to project geometry onto
a planar stress distribution)
requires that:

- 1. We study the three-dimensional nature of faulting.



- a. Importance of higher-order multipolar sources.

b HEALING

Other horrid problems:

- 2. Coupling of dynamical rupture with quasistatic rupture
- 3. Coupling of fluid diffusion with fracture (stress diffusion)
 - ▼ ... and vice-versa.
- a. Fluid flow in conduits of irregular shape
- b. Fluid flow (and stress diffusion) in tortuous geometries



TORTUOSITY

Some Additional Problems & Reminders

1. Statistics: Not only useful to delineate major trends, but also extremely important to give estimates of accuracy of determinations

examples: Seismic Risk Maps, of recurrence times
Extreme value statistics

2. Dilatancy: Where does the extra volume go?
(where does the matter go?)
Role of the surface of the earth.

3. Elasticity: During the process of opening of (diagonal) cracks, the "grains" deform elastically

- 4. What is the nature of friction?
 - ▼ The formation of BOUCHE (mylonite)

5. Based on what we know at this point, can we construct a model in which
 - a. Small egs. occur off (away from) (Geologically) myself far (?)

- b. Large egs. only occur on my own faults

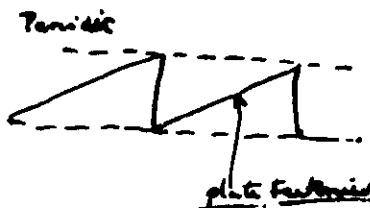
- c. Aftershocks occurs

6. What is influence of healing?
what is nature of healing?

7. What is role of water?

ANY THEORETICAL MODEL
If Earthquakes are MUST SATISFY

perfectly periodic
or (Poisson) random
we do not have to bother to construct complex
models for their occurrence.



If earthquakes are CLUSTERED)

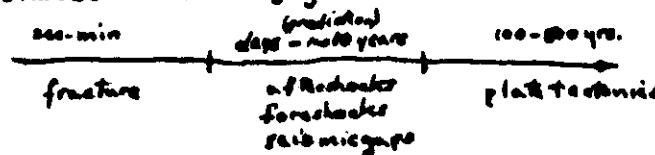
INTERACTION: we must develop an interactive model.

→ Stress redistribution due to fracture

But stresses are redistributed elastically
TOO FAST!

We cannot account for aftershocks

timescales of changing stress



We need a time delay RHEOLOGY

→ Creep (non-elastic) processes giving time delays

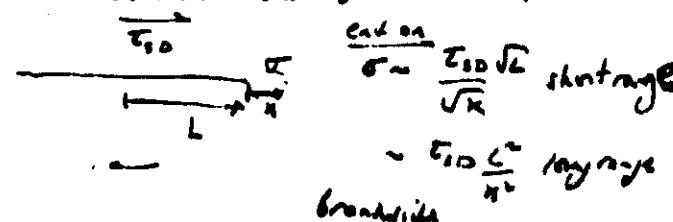
→ plus geometrical constraints
FAULTS!

Large earthquakes (especially) must take into account
Finite size of sources
e.g. $\dots \rightarrow a = b = c \rightarrow a$

CONSTRAINTS THIS FAR

1. PLATE TECTONICS
2. Non-Elastic CREEP RHEOLOGY
3. BRITTLE (ELASTIC) FRACTURE
4. GEOMETRY OF FRACTURE SURFACES & ORIENTATIONS

Summary of Lecture 2

1. Three time scales of seismic phenomena
 - a. To analyse seismograms in a given period range, we should have a good model of the source in that period range
 - b. What causes the time delays in aftershocks?
2. Creep phenomena in laboratory expts.
 - a. Water saturated rocks creep before they break.
 - b. Accelerated creep prior to fracture
 - c. Time delays to fracture decrease with increasing external stress $t_p \sim e^{-(\sigma_p/\sigma) \text{const}}$
3. Increasing fracture (earthquake) size of aftershock prior to Main Shock.
 - a. Explanation of seismic gaps & doglegs
 - b. Stress redistribution upon fracture
 - c. Crack Fusion! $\sigma_m = \frac{\sigma_0}{2}$
4. Anisotropy, multiple fractures in large shock
 - a. Solution in Water of polycrystalline rocks & re-crystallization
5. Low Viscosity zone in lower crust (long-dit trans. oblique)
 - a. Ductility (plasticity) of lower crust: brittle-ductile transition

