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ASPECTS OF HELIOSEISMOLOGY

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Effects of Stratification

Lecture 4

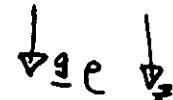
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ASPECTS OF HELIOSEISMOLOGY

The inclusion of gravity introduces a body force into the eqn. of momentum:

$$\rho \left(\frac{\partial \mathbf{x}}{\partial t} + (\mathbf{v} \cdot \nabla) \mathbf{x} \right) = -\nabla p + \underline{j} \times \underline{B} + \underline{g} \rho$$

where $\underline{g} = g \hat{z}$ for \hat{z} pointing downwards
($g = 274 \text{ m s}^{-2}$)



Thus we have: a preferred direction (the z-axis)

an additional force

and imposed length and timescales

- these are additional to any the magnetic field may provide.

Absence of B: Equilibrium

In equilibrium ($\frac{\partial}{\partial t} \approx 0, \nabla \approx 0$)
we have $p = p_0, T = T_0, e = e_0$:

$$\nabla p_0 = +g \rho_0 \hat{z}, \text{ i.e. } p'_0(r) = +g \rho_0. \quad (' \equiv d/dr)$$

Combined with gas law:

$$P = \frac{k_B e}{\mu M_H} T$$

Thus,

$$p'_0 = +g \left(\frac{\mu M_H}{k_B T_0} \right) p_0$$

Boltzmann constant, k_B
 M_H mass of hydrogen atom
 μ mean particle mass

We may integrate to obtain

$$p_0(z) = p_0(0) e^{n(z)}$$

$$\rho_0(z) = \rho_0(0) \frac{\Lambda_0(0)}{\Lambda_0(z)} e^{n(z)}$$

where

$$\Lambda_0(z) = \frac{p_0}{p_0'} = \frac{p_0}{g\sigma} = \frac{k_B T_0(z)}{\mu M_H g} \quad \text{pressure scale-height}$$

$$\text{and } n(z) = \int_0^z \frac{dz}{\Lambda_0(z)}, \quad \text{the 'number' of scale-heights between } z=0 \text{ and } z.$$

($z=0$ is an arbitrary reference level, usually taken to be at optical depth $\tau_{0.5}=1$ in the photosphere).

Examples 1. Isothermal atmosphere ($\Lambda_0 = \text{constant}$, assuming μ constant)

$$\text{Then } \text{Also } n(z) = z/\Lambda_0$$

$$\text{and } p_0(z) = p_0(0) e^{-z/\Lambda_0}, \quad \rho_0(z) = \rho_0(0) e^{-z/\Lambda_0} \quad \text{in height}(z)$$

i.e. pressure and density decline exponentially fast, with a scale of Λ_0 (i.e. pressure + density fall by a factor of $e(=2.718)$ in a height of Λ_0). At the temperature minimum, $\Lambda_0 \approx 100 \text{ km}$

2. Linear temperature (constant temperature gradient)

$$\text{With } \Lambda_0(z) = \Lambda_0(0) + z\Lambda'_0, \quad \Lambda'_0 \text{ constant} \quad (\text{and dimensionless})$$

$$\text{Then } n(z) = \frac{1}{\Lambda'_0} \log_e \left(1 + z \frac{\Lambda'_0}{\Lambda_0(0)} \right)$$

(1')

Introduce

$$z_0 \equiv \frac{\Lambda_0(0)}{\Lambda'_0} = \frac{c_s^2(0)}{(c_s^2)'} \quad \text{temperature scale-height at } z=0$$

for sound speed c_s :

$$c_s^2(z) = \frac{p_0}{\rho_0} = c_s^2(0) + (c_s^2)' z = c_s^2(0) \left(1 + \frac{z}{z_0} \right)$$

and

$$m = \frac{z_0}{\Lambda_0(0)} - 1; \quad m+1 = \frac{\gamma_g}{(c_s^2)'} \quad ,$$

Then

$$\frac{p_0(z)}{p_0(0)} = \left(1 + \frac{z}{z_0} \right)^{m+1}, \quad \frac{\rho_0(z)}{\rho_0(0)} = \left(1 + \frac{z}{z_0} \right)^m,$$

$$\text{so } p_0(z) \propto p_0 \left(1 + \frac{z}{z_0} \right)^{m+1}$$

m is called the polytropic index.

(1'')

General Considerations

With no stratification, no g , we have
 (isotropic) sound waves $c_s = (\gamma \rho_0)^{1/2}$

If $g \neq 0$, then sound waves must be
anisotropic.

Also, because of the new force ($g \neq 0$) another mode of oscillation arises:
gravity waves

These two modes are also called
 - p - and g - modes

Lengthscales The inclusion of g introduces a lengthscale
 $H = \frac{\rho_0}{\rho'_0}$ density scale ht.

Similarly, $\Lambda = \Lambda_0 = \rho_0 / \rho'_0$ pressure scale ht.

Combined with c_s , we get a

Timescale In terms of frequency,

$$\omega_a = \frac{c_s}{2H}$$

acoustic cutoff freq.

This is the cutoff frequency. (Can also construct

$$\omega^2 = \frac{g}{\rho} - g^2$$

(second) buoyancy freq.

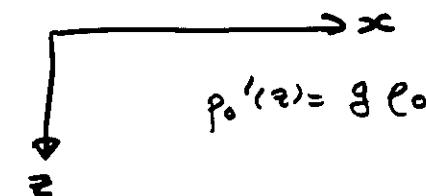
(2)

Sound Waves in Stratified Atmosphere

$$B=0$$

plane atmosphere

treatment actually includes p and g-modes, f-mode
 and unstable g-modes (convection).



$$\rho'_0(z) = g \rho_0(z)$$

$$Eqm \Rightarrow \rho_0(z) = \rho_0(0) \exp\left(\int_0^z \frac{dz}{\Lambda}\right)$$

$$\Lambda(z) = \frac{\rho_0(z)}{\rho_0(0)g} = \frac{k_B T_0(z)}{mg}$$

pressure gradient

$$(Note: c_s^2(z) = \gamma g \Lambda(z))$$

Perturbations

Lamb (1932), Spiegel + Uno (1964)

Introduce $\Delta = \text{div } \mathbf{v}$

Consider [2D] motions $\mathbf{v} = (v_x, 0, v_z) \exp i(\omega t - k_x z)$

Then

$$\left. \begin{aligned} (g^2 k_x^2 - \omega^4) v_z &= \omega^2 c_s^2 \frac{d\Delta}{dz} + g(\gamma \omega^2 - k_x^2 c_s^2) \Delta \\ \frac{dv_z}{dz} + \left(\frac{g k_x^2}{\omega^2}\right) v_z &= \left(1 - \frac{k_x^2 c_s^2}{\omega^2}\right) \Delta \end{aligned} \right\} \quad (*)$$

f-mode Set $\Delta = 0$. Then (*) are satisfied if $\omega^2 = g k$
 (provided b.cndns. on v_z are met).

p-and g-modes

Elim. v_2 :

$$\frac{d^2\Delta}{dz^2} + \left(\frac{c_s^{2z}}{c_s^2} + \frac{\gamma g}{c_s^2} \right) \frac{d\Delta}{dz} + \left\{ \frac{\omega^2 - k_x^2 c_s^2}{c_s^2} - \frac{g k_x^2}{\omega^2} \left(\frac{c_s^{2z}}{c_s^2} - \frac{(\gamma-1)g}{c_s^2} \right) \right\} \Delta = 0 \quad (4)$$

Introduce $\Omega = \rho_0^{1/2} c_s^2 \Delta$ to give

$$\frac{d^2\Omega}{dz^2} + \kappa^2(z)\Omega = 0$$

where

$$\kappa^2(z) = \frac{\omega^2 - \hat{\omega}_a^2}{c_s^2} + k_x^2 \left(\frac{\omega_g^2}{\omega^2} - 1 \right)$$

$$\hat{\omega}_a^2 = \frac{c_s^2}{4H^2} (1 + 2 + l') \quad , \quad H = \frac{\rho_0}{\rho}$$

$$\omega_g^2 = \frac{g}{H} - \frac{g^2}{c_s^2} \quad \text{buoyancy freq.}$$

H, c_s constant. Waves for $k_z^2 > 0$.

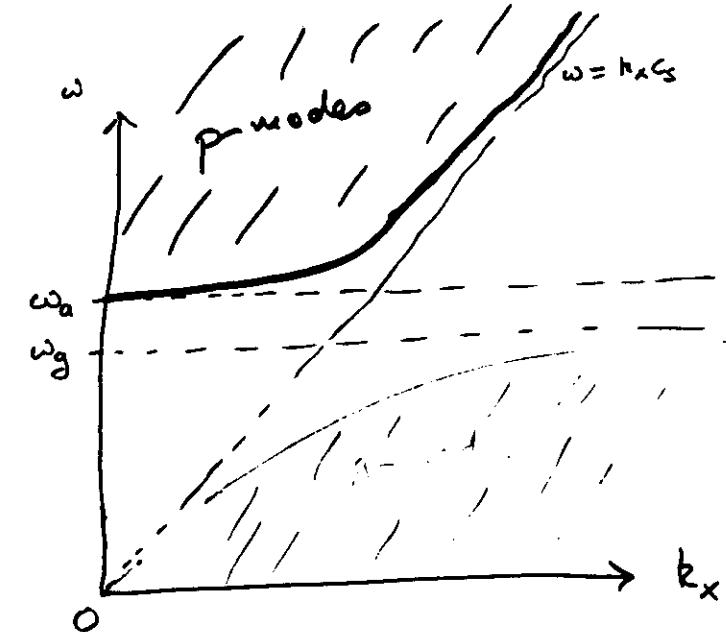
Isothermal atmosphere

Dispersion reln. ($\Omega \propto e^{-ik_z z}$)

$$k_z^2 = \frac{\omega^2 - \hat{\omega}_a^2}{c_s^2} + k_x^2 \left(\frac{\omega_g^2}{\omega^2} - 1 \right)$$

waves for $k_z^2 > 0 \rightarrow$ diagnostic diagram.

(4)



(5)

Diagnostic diagram for Isothermal atmosphere

$$\omega_a^2 = \frac{c_s^2}{4H_0^2}, \quad \omega_g^2 = \frac{g}{H_0} \left(\frac{\gamma-1}{\gamma} \right)$$

solar photosphere at the temp. min.

$$H_0 = 100 \text{ km}, \quad v_a \equiv \frac{\omega_a}{2\pi} = 5.38 \text{ m/s} \quad \}$$

$$\tau_a = 1/v_a = 186 \text{ s}$$

$$v_g \equiv \frac{\omega_g}{2\pi} = 5.27 \text{ m/s} \quad \}$$

$$\tau_g = v_g^{-1} = 190 \text{ s}$$

Linear Profile

$$c_s^2(z) = c_0^2 \left(1 + \frac{z}{z_0}\right), \quad z \geq 0 \quad (6)$$

$$\text{So } p_0(z) \propto \left(1 + \frac{z}{z_0}\right)^{n+1}, \quad q_0(z) \propto \left(1 + \frac{z}{z_0}\right)^n$$

$$\text{where } m+1 = \frac{x_0}{\lambda(x_0)} = \frac{x_0}{(c^2)}, \quad \left[\begin{array}{l} \text{Note: } m \downarrow \text{ if temp. grad} \\ \qquad \qquad \qquad \uparrow \text{ goes up} \end{array} \right]$$

$$\text{Then } \frac{\omega^2}{\omega_0^2} = \frac{m(m+2) \frac{c^2}{r^2}}{4\epsilon_0(r+z_0)}, \quad \omega_0^2 = [m - \frac{1}{r}(m+1)] \frac{g}{(r+z_0)}$$

Eqn. (a) has exact soln. (Lamb)

$$\Delta = e^{-k_x(z+z_0)} \{ C M_{(-a, m+2, 2k_x z + 2k_x z_0)} + D U_C \rightarrow \}$$

m, n confluent hypergeometric fun

C, D arb. constants

$$2a = \left(\frac{m+1}{8}\right) \Omega^2 + \left((\frac{8-1}{8})^m - \frac{1}{8}\right) \frac{1}{\Omega^2} \quad -(m+8)$$

$$J^2 \equiv \omega/g_{k_x}$$

Polytropic Take $\frac{c_0^2}{z_0} = 0 \Rightarrow c_0^2 + \frac{c_0^2}{z_0} \neq 0$

Then if $v_2 = 0$ at $\epsilon = h$ (base)

and v finite at $\varepsilon = 0$ $(\frac{+top}{})$

We get $D=0$ and the dispersion reln. (Spiegel + Umap) \Rightarrow

$$2\sqrt{2} \frac{M'}{M} = 1 + \sqrt{2} - \frac{(m+1)\sqrt{2}}{R_2 h}$$

6

When $\frac{t}{\lambda} h \gg 1$, the paraxial dispersion relation yields

$$a \approx n -$$

$$\therefore [(m+1)\lambda^4 - \gamma(m+2n)\lambda^2 + (\gamma-1)m - 1] = 0$$

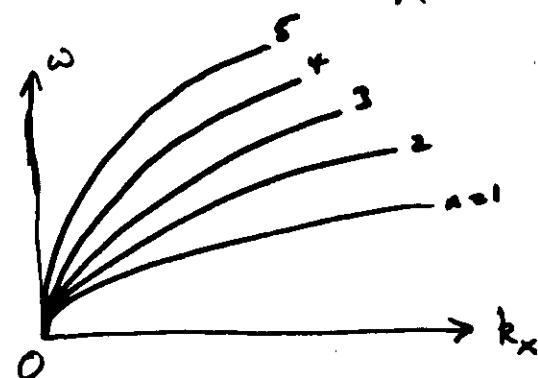
Two roots : p-modes \rightarrow Sturmian sequence of $\omega^2(\beta l^2)$
 g-modes \rightarrow anti-Sturmian " frequency accumulating at $\omega^2 = 0$

If $(\zeta-1) \omega_c < 1$ (i.e. if $\omega_g^2 < 0$), then the g-modes are UNSTABILE.

for adiabatic stratification ($\rho_0 g = 0$), g-modes give $\omega^2 = 0$

and γ -modes are

$$R^2 = R_n^2 \equiv 1 + \frac{2n}{M}$$



parabolas
 $y^2 = kx$

Such modes are observed!

Observations. Parabolas first obtained by Deubner ('75), when no theoretical consensus existed (to explain 5th oscill. reported by Wrighton et al. in 1962); Deubner's measurements showed that the models proposed (indep.) by Ulrich ('70) and Leibacher + Stein ('77) were the correct ones.

$$(m+1)\Omega^4 - \gamma(m+2n)\Omega^2 + [(\gamma-1)m - 1] = 0$$

- quadratic for $\Omega^2 \equiv \omega^2/gk_x$ (Spiegel + Chao '62)

- 2 roots, + $\rightarrow p\text{-modes}$
- $\rightarrow g\text{-modes}$

Large n: $\omega^2 \sim \omega_n^2$ where

$$\omega_n^2 = \begin{cases} 2\gamma g k_x (n + \frac{1}{2}m) / (m+1) & \rightarrow p\text{-modes} \\ 2\gamma g k_x (m - \frac{m+1}{8}) / (m+2n) & \rightarrow g\text{-modes} \end{cases}$$

p-modes are Sturmian, i.e.

$$\omega_1^2 < \omega_2^2 < \omega_3^2 < \dots < \omega_n^2 < \dots$$

with $\omega_n^2 \rightarrow 0$ as $n \rightarrow \infty$

g-modes are anti-Sturmian, i.e. (with $(\gamma-1)m > 1$)

$$\omega_1^2 > \omega_2^2 > \omega_3^2 > \dots > \omega_n^2 > \dots$$

with $\omega_n^2 \rightarrow 0$ as $n \rightarrow \infty$

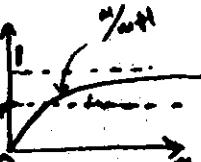
g-modes become unstable if $(\gamma-1)m < 1$

i.e. if $\omega g^2 < 0$

- this corresponds to convection.

Convection occurs if $\frac{m}{m+1} < \frac{m_{\text{adiab}}}{1+m_{\text{adiab}}}$
 $(\gamma = g_2: \approx < \frac{3}{2}) \quad \text{i.e. } \frac{m}{m+1} < \frac{1+m_{\text{adiab}}}{1+m_{\text{adiab}}}$

Schwarzschild (1906)
criterion



Low and intermediate degree modes

Libbrecht + Kaufman (1988, Ap.J. 324, 1172) 8

FREQUENCIES OF HIGH-DEGREE SOLAR OSCILLATIONS

1181

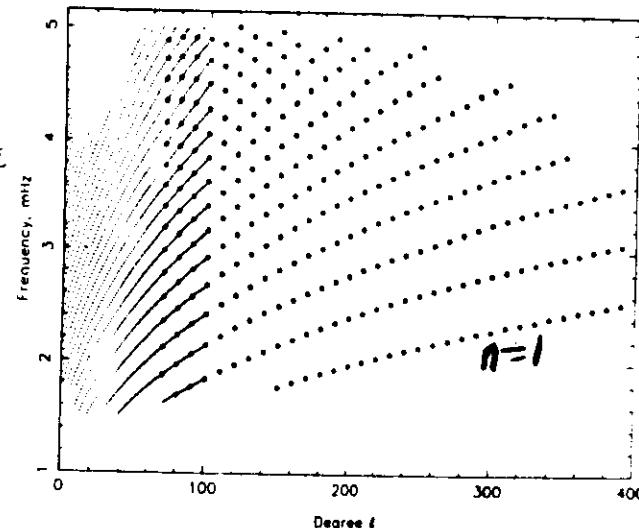


FIG. 5.—Schematic l - v diagram of the p -mode frequencies added by the intermediate-resolution data set (open circles) along with the points from DAI, PP and the no-resolution analysis. The lowest- n ridge has $n \approx 1$. The intermediate-resolution points are listed in Table 2.

High degree modes

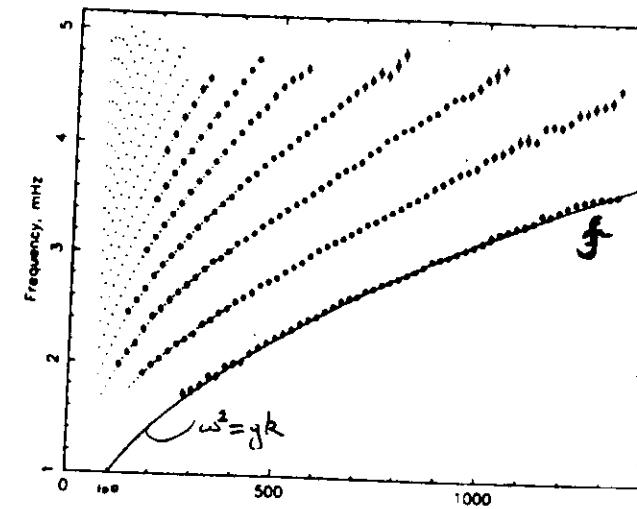
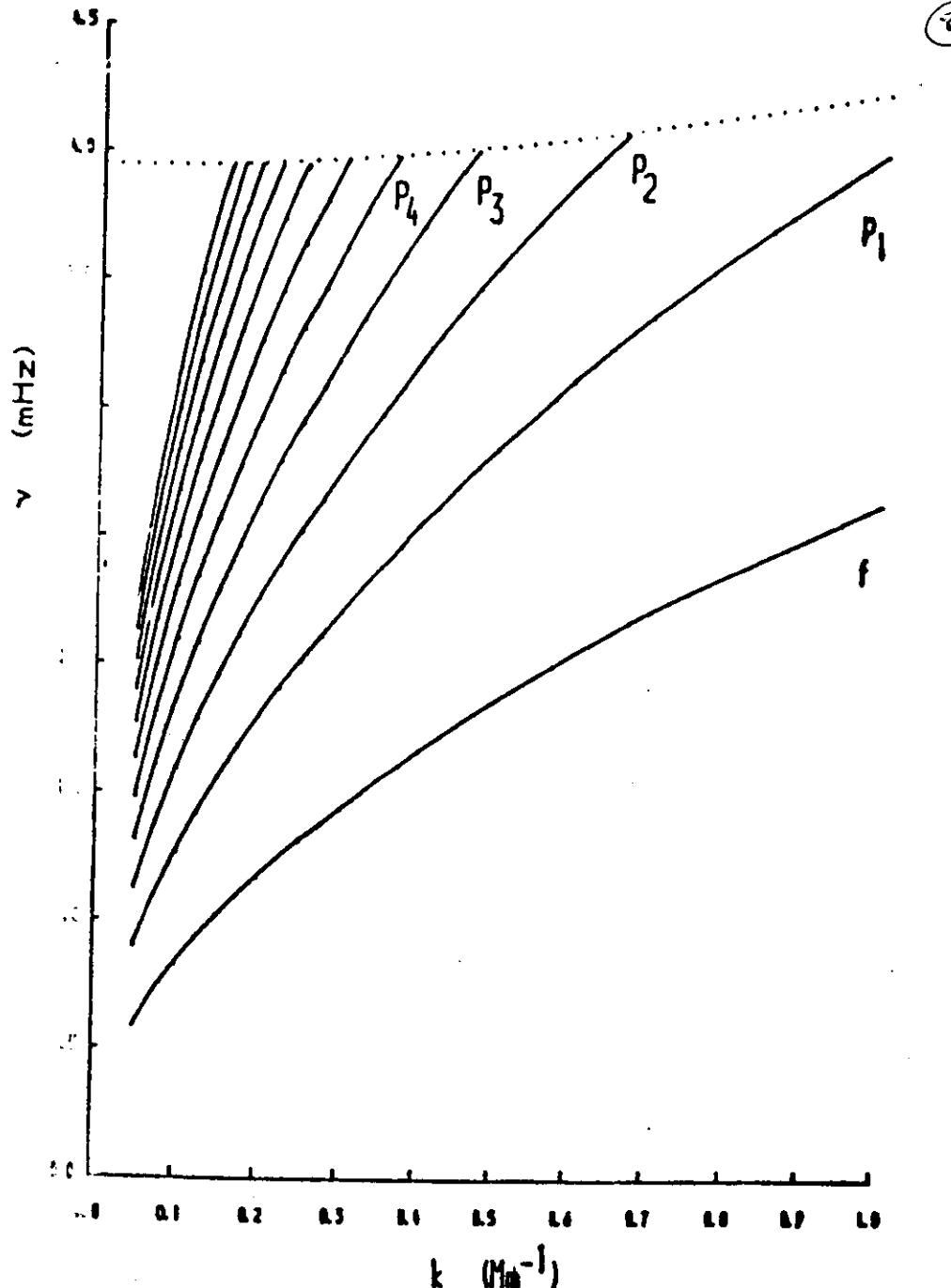


FIG. 6.—Schematic l - v diagram of the p -mode frequencies added by the high-resolution data set (open circles), along with the intermediate-resolution points. The lowest- n ridge has $n \approx 0$, consisting of the f -modes. The smooth curve is the analytic expression for the mode frequencies, $\omega^2 = gk$. After considering residual systematic errors in this data set, the points are not significantly displaced from the line (see text).

Observations



71

90

WKB Aspects

$$\text{Take } \kappa^2 = \frac{\omega^2}{\epsilon_0(\epsilon)} - k_x^2$$

(setting $\omega_0^2 = 0$ and neglecting $\hat{\omega}_n^2$)

Bohr-Sommerfeld reln. gives

$$\int_{\text{cavity}} k dz \sim (n+\alpha)\pi$$

$\frac{xx}{xx} (n \rightarrow \infty)$

phase α ($= 1/2$ in 2 point cavity)

E.g. $\frac{\text{Polytropes}}{\text{Take top of cavity as } z=0 \text{ and bottom at } \kappa^2(z_c)=0}$, viz.

$$z = z_c = (n+1) \frac{\omega^2}{\epsilon_0 k_x^2}$$

$$\therefore \int_0^{z_c} \left[\frac{\omega^2(n+1)}{\epsilon_0 z} - k_x^2 \right]^{1/2} dz \sim (n+\alpha)\pi$$

$$\Rightarrow \frac{\omega^2}{\epsilon_0 k_x^2} = \frac{2}{m} (\alpha + n) = 1 + \frac{2n}{m} \quad (\text{as before})$$

provided $n = n_{1/2}$

and then

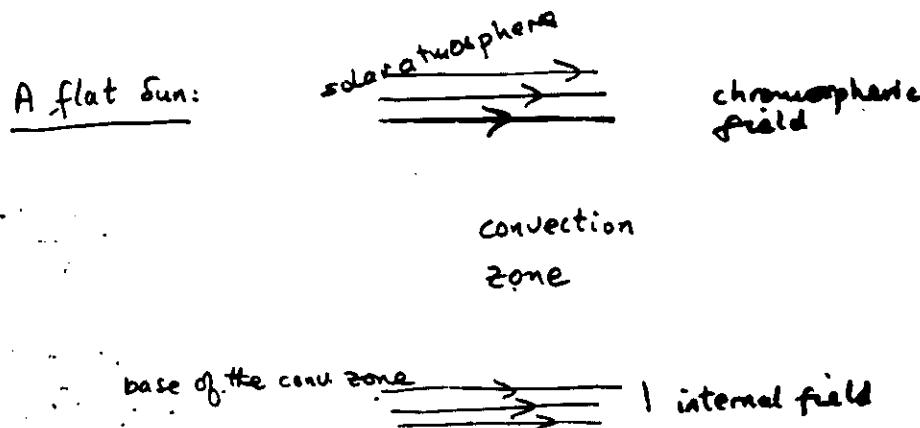
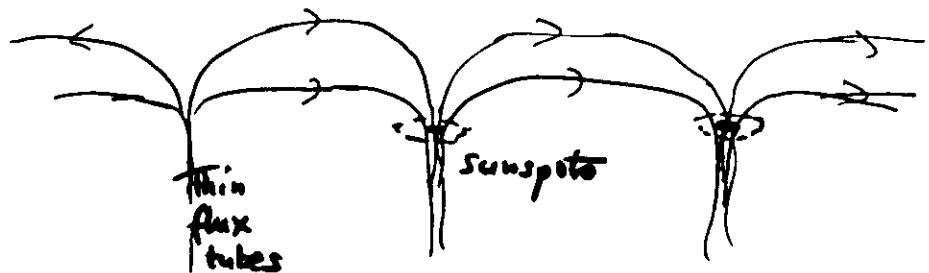
$$\text{cavity depth } z_c = \frac{2(n+\alpha)}{k_x} \propto \frac{\text{wavelength}}{k_x}$$

Eqn. $\frac{xx}{xx}$ may be cast in the form of Abel's integral eqn. and inverted [with $k_x^2 = \frac{2(2+\alpha)}{r^2}$ for a sphere] to get the sound speed. This has been done for the sun (Christensen-Dalsgaard et al. '85).

Thus, helioseismology has given us the map of sound speed in the sun

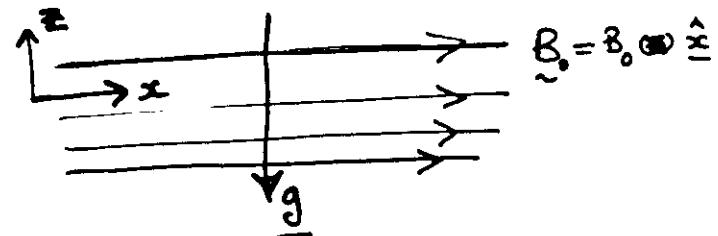
MAGNETIC EFFECTS

Magnetic Canopy



(10)

Effect of a horizontal B



vertical stratification
in $C_S(z)$.

$$\text{EQM. } \left(\rho_0(z) + \frac{B_0^2(z)}{2\rho_0} \right)' = -\rho_0(z) g$$

Linear perturbations: consider [2D] motions,
 $v = (v_x, 0, v_z) e^{i(\omega t - k z)}$

Find (from ideal MHD eqns)

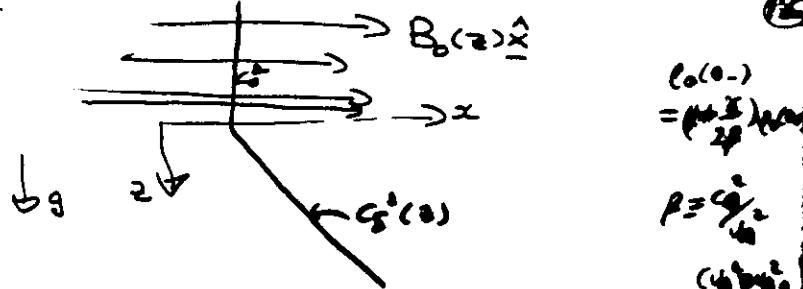
$$0 = \frac{d}{dz} \left\{ \frac{\rho_0(c_s^2 + v_A^2)(\omega^2 - k^2 c_s^2)}{(\omega^2 - k^2 c_s^2)} \frac{dv_z}{dz} \right\} + \\ + \left\{ \rho(\omega^2 - k^2 v_A^2) - \frac{g^2 k^2 \rho}{(\omega^2 - k^2 c_s^2)} - \rho k^2 \left(\frac{\rho c_s^2}{\omega^2 - k^2 c_s^2} \right)' \right\} v_z$$

Complicated eqn! No exact solns. except in very special cases
(e.g. constant c_s, v_A)

Continuous spectra - associated with slow wave

→ includes: fast waves, slow waves, gravity waves,
magnetohydrodynamics, compressible magnetohydrodynamics in
horiz. B , etc. !

Atmosphere



Atmosphere ($z < 0$) is taken to be isothermal.

MHD eqn. is

$$\frac{d}{dz} \left\{ \frac{\rho_0(c_s^2 + v_A^2)(\omega^2 - k^2 c_s^2)}{(\omega^2 - k^2 c_s^2)} \frac{du_0}{dz} \right\} = \left\{ \frac{\rho_0 g^2 k^2}{\omega^2 - k^2 c_s^2} - \rho_0 (\omega^2 - k^2 c_s^2) \right. \\ \left. - g k^2 \left(\frac{\rho_0 g^2}{\omega^2 - k^2 c_s^2} \right)^{1/2} \right\} u_0$$

Take $v_A = \text{constant}$. Then

$$u_z = u_0 e^{z/2H_B} (-1 \pm \sqrt{1 - 4AH_B^2}) \quad , \quad z < 0$$

$$H_B^{-1} = \frac{\rho_0'}{\rho_0} = \frac{\rho_g}{c_s^2}, \quad \Gamma = \frac{\sigma}{(1 + \frac{\sigma}{2\rho})} \quad [R = \sigma, \quad \chi_B = H_B \text{ if } R = 1]$$

Suppose $4AH_B^2 < 1$, and choose + sign so $e^{z/2H_B} u_0 \rightarrow 0 \text{ as } z \rightarrow -\infty$.

Across $z=0$: $u_z \rightarrow 0$

$$\frac{\rho_0(c_s^2 + v_A^2)(\omega^2 - k^2 c_s^2)}{(\omega^2 - k^2 c_s^2)} \frac{du_0}{dz} + \left(\frac{gk^2 \rho_0 c_s^2}{\omega^2 - k^2 c_s^2} \right) u_z \rightarrow 0$$

$\rho_0 u_z^2$ bounded as $z \rightarrow \infty$

In non-magn. region: $\Delta = e^{-k(z+z_0)} U_{(-a, n+2, 2kz + 2kz_0)}, \quad z > 0$
satisfying $\rho_0 u_z^2$ bounded as $z \rightarrow \infty$

Dispersion Reln

Campbell & R. (1989)

Ap.J.

$$2k\omega^2 c_0^2 \frac{U'}{U} + \gamma g \omega^2 - k \omega^2 c_0^2 - g k^2 c_0^2 \\ = \frac{(\omega^2 - k^2 c_0^2)(g^2 k^2 - \omega^4)(c_0^2 + \frac{1}{2} \gamma v_A^2)}{g k^2 c_0^2 + (c_0^2 + v_A^2)(\omega^2 - k^2 c_0^2) \Delta}$$

$$\Delta = \frac{1}{2H_B} [-1 + (1 - 4AH_B^2)^{1/2}]$$

$$U' \text{ eval. at } 2kz_0, \quad z_0 = \frac{c_s^2(0)}{(c_s^2)'}$$

Disp. reln. includes p-modes (modified by B)
f-mode (modif. by B)

MAGNETOACOUSTIC CUTOFF

MHD surface waves

For $k z_0 \ll 1$ may show that

$$v_n = v_0 + C_n(c_s, v_A) k^{m+5/2}$$

↑ freq. in absence
of an atmosphere

← correction due to
an atmosphere

Similar for f-mode:

$$v_n = \bar{v}_0 \left[1 + \frac{(2c_0^2 + v_A^2)v_A^2}{(c_0^2 + v_A^2)c_0^2} \frac{2^{m-1} (m+1)^{m+2}}{m \Gamma(m+2)} (k H_0)^{m+2} \right]$$

Here atmoph. correction is due entirely to the atmosphere
being magnetic. $v \uparrow$ when $B \uparrow$

Atmospheric corrections

$B_0 = 0$

	P_n -modes			
	$n=1$	$n=2$	$n=3$	$n=10$
100	-0.01	-0.009	-0.008	-0.008
500	-8.2	-5.4	-4.8	-5.1
900	-86	-57	-50.5	-54
1300	-375	-248	-220	-236

Atmospheric corrections (to mode P_n) in μHz

$B_0 = 10G$

	$n=1$	2	3	10
100	-0.01	-0.009	-0.008	-0.008
500	-8.2	-5.4	-4.8	-5.2
900	-86	-57	-51	-54
1300	-374	-249	-221	-236

$B_0 = 10^2 G$

	$n=1$	2	3	10
100	-0.01	-0.01	-0.01	-0.01
500	-6.7	-7.2	-7.3	-8.7
900	-70	-76	-76	-92
1300	-306	-329	-321	-399

$B_0 = 200G$

	$n=1$	2	3	10
100	-0.01	-0.02	-0.02	-0.03
500	-4.1	-13.1	-14.9	-19.7
900	-43	-137	-156	-207

(analytical expansion)
 $(\delta = \frac{g}{k}, n = 3/2)$

(13')

Evolution of B_0
(Campbell & R. '89)

$$\Delta v = v(B_0) - v(B_0=0)$$

$$= \text{atmosph. corr.}(B_0)$$

$$- \text{atmosph. corr.}(B_0=0)$$

Low to moderate L : effect negligible.

	$B_0 = 10G$	100G	200G
P_1 :	0.02	1.5	4.1
500	0.2	15.7	43.
900	0.8	68.	187.
1300	-0.02	-1.8	-7.6
P_2 :	-0.2	-25.	-80.
500	-0.8	-83.	-349.
900	-0.03	-3.6	-14.6
1300	-0.4	-38.	-153.0
P_{10} :	-1.6	-164.	-665.7

p-modes

'quiet' sun

Note: negligible
diff. from
 $B_0 = 0$
case

'Active' Sun

$$\Delta v = v(B_0) - v(0)$$

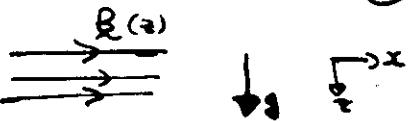
	$B_0 = 10G$	$10^2 G$	200G
500	0.01	0.85	2.79
900	0.10	8.90	29.2
1300	0.44	38.7	127.1

f-mode

$$v(0) = \frac{(gk)^{1/2}}{2\pi}$$

Note: $\Delta v_f > 0$ whereas $\Delta v_{P_n} < 0$ ($n=2, 3, \dots$)

A Magnetic Lamb's Eqn.



Horizontal magnetic field, $B(z)$:

Magnetostatic equil. :

$$\frac{d}{dz} \left(\rho + \frac{B^2}{2\mu_0} \right) = g e \quad \left. \begin{array}{l} \\ \end{array} \right\}$$

$$\rho = \frac{k_B}{m} e T$$

amb. $T(z)$

Perturbations

$$\underline{v} = (v_x(z), 0, v_z(z)) e^{i(\omega t - k z)}$$

Eqns. of ideal MHD yield

$$\left(1 - \frac{k^2 c_s^2}{\omega^2}\right) \Delta = \frac{dv_z}{dz} + \left(\frac{g k^2}{\omega^2}\right) v_z$$

$$c_s^2 \frac{d\Delta}{dz} + (\gamma - 1) g \Delta - \frac{\gamma (BB')}{\mu_0 \rho} \Delta = -g \frac{dv_0}{dz} + (k^2 v_A^2 - \omega^2) v_0$$

$$-\frac{1}{\rho} \frac{d}{dz} \left\{ \frac{B^2}{\mu_0} \frac{dv_0}{dz} \right\}$$

$\Delta \equiv \text{div } \underline{v}$

Elin. of $\Delta \Rightarrow$

$$\frac{d}{dz} \left\{ \rho \frac{(c_s^2 + v_A^2)(\omega^2 - k^2 c_s^2)}{(\omega^2 - k^2 c_s^2)} \frac{dv_0}{dz} \right\}$$

$$= \left\{ \frac{\rho g^2 k^2}{\omega^2 - k^2 c_s^2} - \rho (\omega^2 - k^2 v_A^2) - g k^2 \left(\frac{\rho c_s^2}{\omega^2 - k^2 c_s^2} \right)' \right\} v_z$$

(14)

Compare with Lamb's eqn. when $v_A = 0$:

$$\boxed{\frac{d^2 Q}{dz^2} + \kappa^2(\omega) Q = 0, \quad Q = \rho^k c_s^2 \Delta}$$

where

$$\kappa^2 = \frac{\omega^2 - \omega_b^2}{c_s^2} + k^2 \left(\frac{\omega_g^2}{\omega^2} - 1 \right)$$

Note simplicity!

$$\omega_g^2 = \frac{g}{H} - \frac{g^2}{c_s^2} \quad \text{buoyancy freq.}$$

$$\omega_b^2 = \frac{c_s^2}{4H^2} (1 + 2H') \quad \text{cutoff freq.}$$

$$H = e/e' \quad \text{density scale-height.}$$

What happens if $v_A \neq 0$?

In general, κ is a scalar!

But for $\omega \gg k v_A$ we may obtain (Roberts+Campbell '86)
(constant κ)

$$\kappa^2 = -k^2 + \frac{\omega^2}{c_s^2 + v_A^2} - \frac{\omega_b^2}{c_s^2} + \frac{k^2 \omega_g^2}{\omega^2} \left(\frac{c_s^2}{g^2 + v_A^2} \right) + \frac{k^2 c_s^2 v_A^2}{(c_s^2 + v_A^2)^2} + O\left(\frac{k^2}{\omega}\right)$$

This is the desired expression for a "magnetic Lamb's eqn."

Magnetic effects

p-modes

p-modes scatter off

g-modes - cts. spectrum (destroys anti-symmetric)
 g-mode freq. at low ω_B^2 ($\text{high } n$)

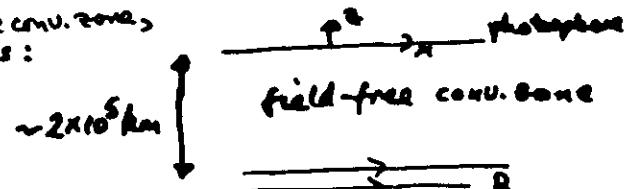
isolated flux tubes
 (Bogdan + Scoville '85)

- sunspots (which form a
 sink for p-modes;
 Braun et al. '87 preprint)
 Andretta and Thomas '87 "

What is B at base of convection zone?

Also, p(modes) interact with B in chromosphere,

and B at base of convection zone. Roberts + Campbell '86,
 in attempt to deduce B at base conv. zone,
 have argued as follows:



The magnetic equivalent of $\textcircled{1}$ may be derived - but is very complicated in algebra. For high frequency limit appropriate for p-modes, R+C get (for $\omega \gg \omega_B$)

$$k^2(\epsilon) \approx -k^2 + \frac{\omega^2}{\xi^2 + \xi^2} = \frac{\omega_0^2}{\xi^2} + \frac{k^2 c_s^2}{\xi^2 + k^2} + \frac{k^2 \eta^2}{\omega^2} \left(\frac{\omega^2}{\omega_B^2} \right)$$

- result is exact for $g=0$, $k=0$, $\omega_B=0$.

Applying standard WKB theory then gives

$$\frac{\omega}{B} = \left(n + \frac{1}{2} + \epsilon \right) (\Delta \nu_0) \left(1 + \frac{\omega_A^2}{2 \xi^2} \right)$$

phase factor

for cyclic freq. $\nu_B \approx \omega/2\pi$, and $k^2 \equiv \frac{2(2+1)}{R_\odot^2}$, $\Delta \nu_0 = \left(\frac{1}{2} \right)^{\frac{1}{2}}$

Consider now the change in freq. due to a change in B over solar cycle. If ν_B is freq. in a field of strength B_0 and $\nu_{B'}$ is freq. in field of strength B' , then (from $\textcircled{2}$)

$$B^2 - B'^2 = 8\pi \rho c_s^2 \frac{1}{\nu_0} (\nu_B - \nu_{B'})$$

where ν_0 = freq. in absence of B ($\nu_0 \approx 3.1 \text{ Hz}$)

Now SMM ACRIM OBSERVATIONS by Woodard and Noyes (1985) give

a freq. change (decrease) of $\approx 0.4 \mu\text{Hz}$ (low n)

from 1980 (near solar max.) to 1984 (approaching min.)

(with $\rho = 0.2 \text{ g cm}^{-3}$, $c_s = 2 \times 10^5 \text{ cm s}^{-1}$)

Applied to the above formulae, gives

B must be at least $5 \times 10^5 \text{ G}$ (This is strong; best dyno calc. estimate $B = 10^4 \text{ G}$)

if a changing B is the cause of a change in freq.

But note: - obsns. are at limit of instrument so uncertain

- Elsworth et al ('87) do not confirm W+N result (ground-based network)

- Very recently, Fassat et al. '87 confirm W+N result on analyzing 704 hrs. of South Pole data; and Woodard has given ACRIM results for '87 - showing them up.

- only a longer chain of obs. (say to 1990) will confirm or refute the claim that p-modes change over the solar cycle.

