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CONTENTS

PART 1

Modelling of Environmental and Climatic Problems: Wind and Water Erosion, Mathematical Modelling & Applications

PART II

Climatological Changing Effects on Wind, Precipitation, Erosivity, Aridity, Large, Meso-and Small Scale Analysis

PART I

MODELLING OF ENVIRONMENTAL AND CLIMATIC PROBLEMS: WIND AND WATER EROSION

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ABSTRACT: Magnitude of wind and water erosion mainly depend on wind velocity, rainfall rate, slope and soil characteristics. The main purpose of this lecture is to define the role of small, meso and large scale phenomena (local and synoptic fluctuations) on water and wind erosion These lecture notes presents some results on wind speed simulation and seasonal fuluctuations of water defict for the selected station in different erosion risque and transition regions of Turkey.

Key Words: Wind speed simulation, wind erosion, water erosion, climatological

1. INTRODUCTION

In order to simulate the process of soil erosion by wind, wind speed and it's characteristics are considered, (Skidmore, 1986, 1995, 2000a, 2000b). Prediction of wind speed and direction is extremely difficult. Raindrops, water and soil type play an important role on soil erosion, flood, landslides etc. The reason of soil erosion is mainly caused by the impact of raindrops on the soil surface and its flow between rills and in channels down slope. It also causes landslides on steep slopes. The erosivity effect of raindrops depends on the energy of a rainstorm, Pla Sentis (1998), Flanagen and Livinston (1995) and Gabriels (1993). Monthly rainfall data were used to compute rainfall erosivity indices for various stations in Ghana, (Oduro-Afriyie, 1996). Temporal and spatial variation of rainfall erositivity and climatic factor of wind erosion are investigated in Turkey, Aslan, (1997), Tulunay et al., (2002), Aslan et al., (2002). Higher values of Fournier Index are observed in the North-eastern Black Sea, the Southern Aegean Sea and Mediterranean Sea Regions. This lecture notes also presents some results on wind and water erosion characteristics at the Northwestern part of Turkey (Gökçeada).

2. METHODOLOGY

2.1. Rainfall Erosivity Factor and Water Erosion

Rainfall and runoff erosivity factors are defined by considering the results of field measurements. The Fournier Index described as a climatic index is defined by Odura-Afriye (1996) as:

 $C_{p} = P_{\max}^{2} / P \tag{1}$

Where; C_p ; Fournier Index (mm), P_{max} ; The rainfall amount in the wettest month and P; Annual precipitation (mm).

Table 1 shows classes of rainfall erosion risk based on the Rainfall Erosivity Index. C_p values above 60 show severe to extremely severe erosion risk in average climatic conditions, Odura-Afriye (1996).

Class No	Erosion Risk Class	Fournier Index	Soil Loss (T/ha year)		
		(C _p)			
1	Very Low	<20	<5		
2	Low	21-40	5-12		
3	Moderate	41-60	12-50		
4	Severe	61-80	50-100		
5	Very Severe	81-100	100-200		
6	Extremely Severe	>100	>200		

Table 1. Classes of rainfall erosion risk, indexes and soil losses

2.2. Soil Moisture Prediction

Soil moisture over the western Turkey (in İstanbul) basin has been evaluated for long-term data by using De Martonne's Index (I) (Piervitali et al., 1999). The index is given by the following equation:

$$I = P / (T + 10)$$
 (2)

Where: P; total yearly precipitation (mm) and t; mean yearly temperature (°C). Index values more than 30, correspond to the humid areas where time adjusted irrigation was necessary.

2.3. Wind Erosion (Climatic Factor)

Wind erosion index is also defined as climatic factor (C_w) . It is a function of horizontal wind speed as given below:

$$C_w \alpha V^3$$
 (3)

Erosion risk classes based on mean wind speed values have been studied by Aslan (1997) and Aslan et al., (2002).

2.3.1. Determination of Daily Wind Speed

The cumulative Weibull distribution function F(u) and probability density function f(u) are defined by Skidmore (1986) and (1995) as below:

$$F(u) = 1 - \exp[-(u/c)^{k}]$$
(4)

Where u is wind speed, k is shape parameter (dimensionless) and c is scale parameter (m/s).

$$F(u) = dF(u)/du = (k/c)(u/c)^{k-1} \exp[-(u/c)^{k}]$$
(5)

$$F_{1}(u) = [(F(u) - F_{o}) / (1 - F_{o})] = 1 - \exp[-(u/c)^{k}]$$
(6)

Where $F_1(u)$ is the cumulative distribution with the calm periods eliminated, and F_0 is the frequency of the calm periods. Wind speed, u, the dependent variable:

$$u = c \{ -\ln[1 - (F(u) - F_o) / (1 - F_o)] \}^{1/k}$$
(7)

The programme draws a random number, 0.0 < RN < 1.0. It is assigned to F(u), and subtracted from it the frequency of calm periods, F_0 .

2.3.2. Determination of Sub-daily Wind Speed

Program reads from the wind data-base the ratio of maximum to minimum mean hourly wind speed and the hour of maximum wind speed for the location and month under consideration. Calculate the maximum and minimum wind speed for the day based on the representative wind speed as calculated above and given the ratio of maximum to minimum wind speed:

$$urep = (umax + umin) / 2$$

$$uratio = umax / umin$$
(8)
(9)

where, urep is the daily mean representative wind speed as calculased from Eq. 2.5, uratio is the ratio of daily maximum (umax) to daily minimum (umin) wind speed. Wind speed for any hour of the day u(I) can be simulated from:

$$u(I) = urep + 0.5 (umax - umin) cos [2\pi (24 - hrmax + I)/24]$$
 (10)

where hrmax is the hour of the day when wind speed is maximum; I is index for hour of day and the other variables are as previosly defined.

2.4. Aridity Index

Aridity index (AI) is given by the following equation, Türkeş (1999) Aslan and Tokgözlü (2000).

$$AI = P/PE$$
(11)

where, P; Annual total precipitation (mm) and PE; Potential evaporation (mm). Aridity Index values for arid and dry sub-humid areas have been changed between 0.05 and 0.65.

3. ANALYSIS

3.1. Analysis of Rain-Erosivity

Time variation of regional average annual total precipitation values in Turkey shows a increasing trend between 1900-1998. Erosivity values determined for overall over Turkey (average value) show severe erosion risk in winter, (Table 2). It gives some results on statistical characteristics of annual regional total precipitation and FI values based on the "Climate, Impacts LINK Project" (Giorgi and Francisco, 2000; New et al, 1999, 2000), Aslan et al., (2002).

Year	Spring	Erosion Class	Summer	Erosion Risk Class	Autum n	Erosion Class	Winter	Erosion Class	Mean	Erosion Class
1900-1930	20,9	Very Low	10,4	Very Low	54,2	Moderate	78,2	Severe	40,9	Low
1931-1960	30,7	Low	28,5	Low	47,5	Moderate	59,8	Moderate	41,6	Moderate
1961-1998	36,3	Low	24,5	Low	47,0	Moderate	60,7	Severe	42,1	Moderate
1900-1998	33,9	Low	27,1	Low	48,2	Moderate	59,6	Moderate	42,2	Moderate

Table 2. Seasonal variation of erosivity (Fournier Index) and erosion risk class in Turkey

3.2. Analaysis of Wind Speed and Erosion

Data used in this study is hourly wind speed measurements by using automatic wind recording system which was mounted in Gökçeada (Tuzburnu, Altitude: 34m msl, Latitude:40° 11'N, Longtitude: 25° 54'E) between 1997 and 1998 and , Ugurlu, Çinaralti and National Station, between 1992-1993, 1997-1998). To define water erosion at the study area, monthly and annual rainfall rate values based on long term observations are analysed.

When the other wind speed values are considered at Ugurlu, Çinaralti and National Station between 1992-1993, the linear regression coefficient (r) between wind speed observations (u) and theoretical values is defined as 0.97 with the significant level 1 (confidence limits 0.99). The linear regression coefficient (r) between wind speed observations (u³) and theoretical values at all stations in Gökçeada (Tuzla, Ugurlu, Çinaralti and National Station, between 1992-1993, 1997-1998), is 0.94 with the significant level 1 (confidence limits 0.99), respectively, (Figure 1).

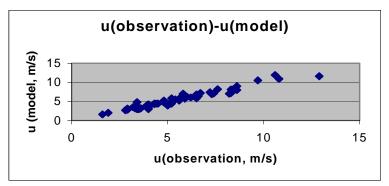


Figure 1- Linear Relation Between Average Wind Speed Values at Gökçeada (Tuzla, Uğurlu, Çinaralti, National Station, 1992-1993, 1-Sept.1997-10-August, 1998).

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REFERENCES

Aslan, Z, (1997)- Analysis of Rainfall Erosivity and Climatic Factor of Wind Erosion. ICTP Pre-print, IC/97/135, Trieste, Italy.

Aslan, Z. and A.Siddiqi, (1999): "Modelling of Environmental and Climatic Problems: Prediction of Wind and Water Erosion", ICIAM99, MSP-243, 4-9 July, Edinburgh.

Aslan, Z., and A. Tokgözlü, (2000)- Climatological Changing Effects on rainfall Erositivty and Wind Erosion", AGROENVIRON2000: 2nd. International Symposium on New Technologies for Environmental Monitoring and Agro-Applications, Tekirdağ, Turkey, pp.265-273.

Aslan, Z., E. Skidmore, E. Feoli, D. Maktav, H. Erol, F. S. Erbek, D. Okçu, A. S. Sogut, P. Giacomich, S. Mauro, K. Mighozzi, (2002), "The Use of Conventional Data and Remote Sensing for Classification of Erosion and Land Degradation", 3rd. AGROENVIRON Symposium, 26-28 October, Cairo.

Flanagen, D.C and J. Livingston, (1995)- USDA-Water Erosion Prediction Project (WEPP), Soil and Water Conservation Society, Iowa, USA, pp.131.

Gabriels, D., (1993)- The USLE for Predicting Rainfall Erosion Losses. ICTP-SMR. 705-3. Trieste, Italy.

Giorgi, F., and R. Francisco, 2000- Evaluating uncertainties in prediction of regional climate changing. Geophysical letters, Vol. 27, No. 9, pp. 1295-1298.

New, M. G., M. Hulme and P. D. Jones, (1999)- Representing 20th century space –time climate variability. I: Development of a 1961-1990 mean monthly terrestrial climatology. J. Climate. Vol. 12, pp. 829-856.

New, M. G., M. Hulme and P. D. Jones, (2000)- Representing 20th. Century space- time climate variability. II: Development of 1901-1996 monthly terrestrial climate fields. J. Climate. Vol. 13, pp. 2217-2238.

Oduro-Afriye, K., (1996)- Rainfall Erosivity Map for Ghana", Geoderma, Elsevier Science B.V., 1125, pp.6.

Piervitali, E., M. Conte and M. Colacino. (1999)- Rainfall Over the Central-Western Mediterranean Basin in the Period, 1951-1995.II. Precipitation Scenarios. Nuovo Cimento C. 22C. 5. pp. 649 - 661.

Pla Sentis, I., (1998)- Modeling the Influence of Soil Sealing and Soil Compaction on Soil Erosion Processes, College on Soil Physics, ICTP/SMR.1065-9, Trieste, Italy.

Skidmore, E. L., (1986)- Wind Erosion Climatic Erosivity, Climate Change, 9, pp. 195-208.

Skidmore, E.L., (1995)- Wind Erosion Climatic Erosivity, ICTP College on Soil Physics, SMR., 873-19, Trieste, Italy.

Skidmore, E., L., (2000a)- Air, Soil, and Water Quality as Influenced by Wind Erosion and Strategies for Mitigation,AGROENVIRON2000:2nd.International Symposium on New Technologies for Environmental Monitoring and Agro-Applications ",Tekirdağ, pp.216-221.

Skidmore, E., L., (2000b)- Sustainable Agriculture: Actions and Strategies, AGROENVIRON2000: 2nd. International Symposium on New Technologies for Environmental Monitoring and Agro-Applications", Proceedings (Workshops), Tekirdağ, Turkey, pp. 1-3.

Tulunay, E., E. T. Senalp, Y. Tulunay and Z. Aslan, (2002): "Development of Neural Net Based Models for Non-Linear Agro-Environmental systems", 3rd. AGROENVIRON Symposium, 26-28 October, Cairo.

Turkes, M., (1999)- Vulnerability of Turkey to Desertification With Respect to Precipitation and Aridity Conditions. Tr. J. of Engineering and Environmental Science. 23. pp. 363-380.