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"7th Workshop on Three-Dimensional Modelling
of Seismic Waves Generation and their Propagation"

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**Study of Seismic Swarms Associated
with Fluid Intrusions**

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Study of seismic swarms associated with fluid intrusions

ICTP Course 2004 Trieste

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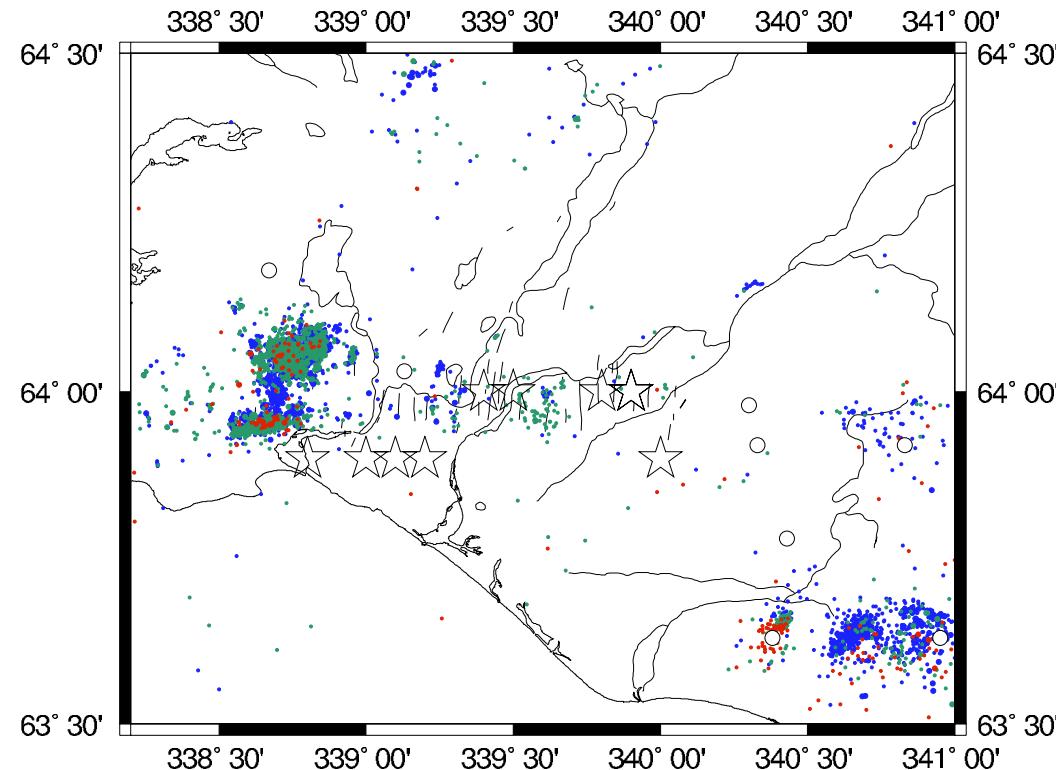
Content

1. Physics of fluid intrusions and earthquake swarm models
2. study of seismic swarms
3. Practical: relative moment tensor inversion

fluid-induced swarms - observations

- Typically in volcanic / geothermal regions
- long activity with possible key events or bursts, $M < 4$
- large b-values ($b > 1$), e.g. up to 2.5
- no typical Omori law decay
- possible hypocenter migration
- often large double-couple components, partly positive isotropic component, rarely pure tensile events
- accompanied by long-period or tremor events, ground deformation or degassing

South Iceland Seismic Zone



○ volcanos

☆ $M \geq 6$ earthquakes since 1700

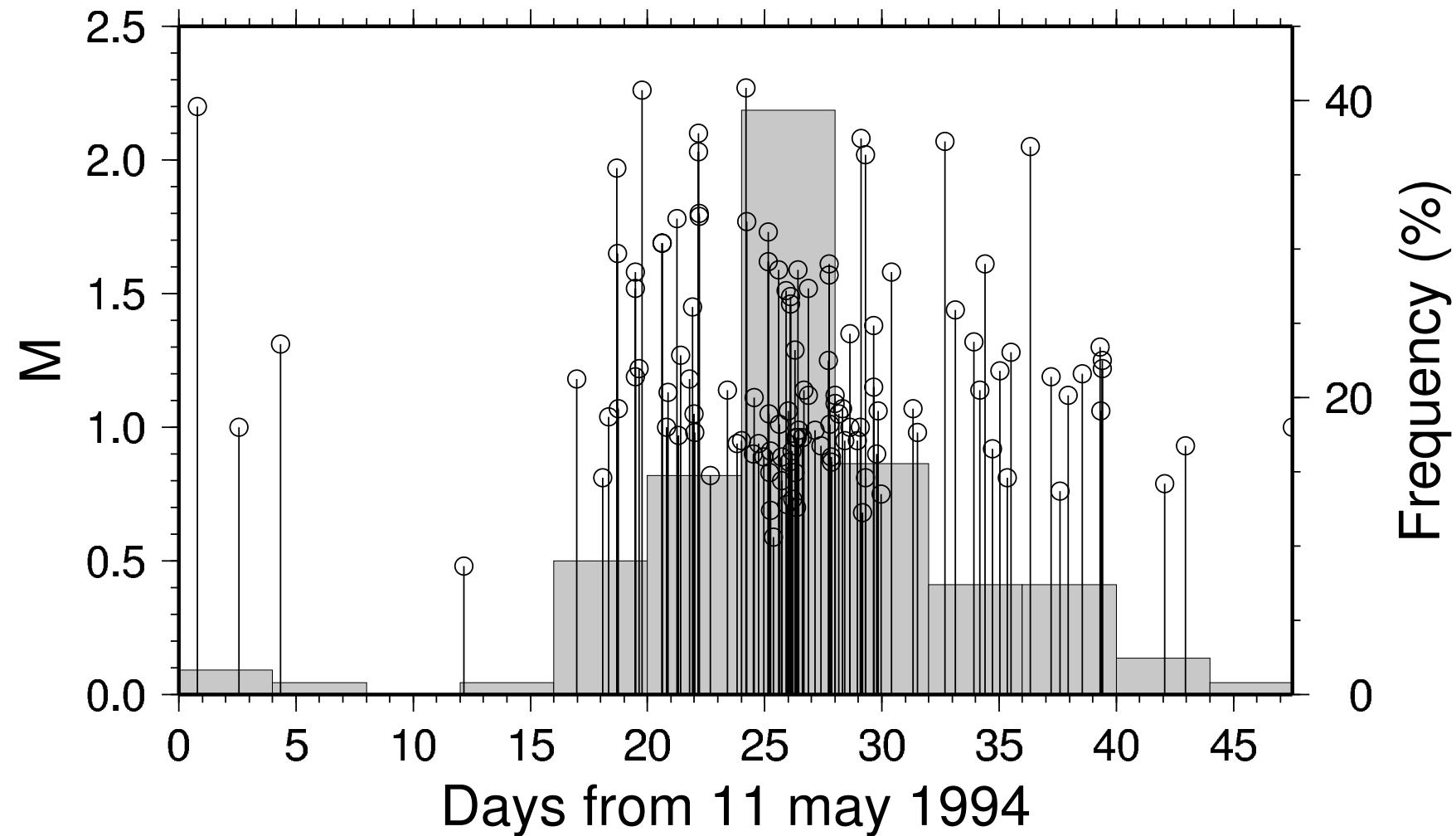
○ 0 - 5 km

○ 5 - 10 km

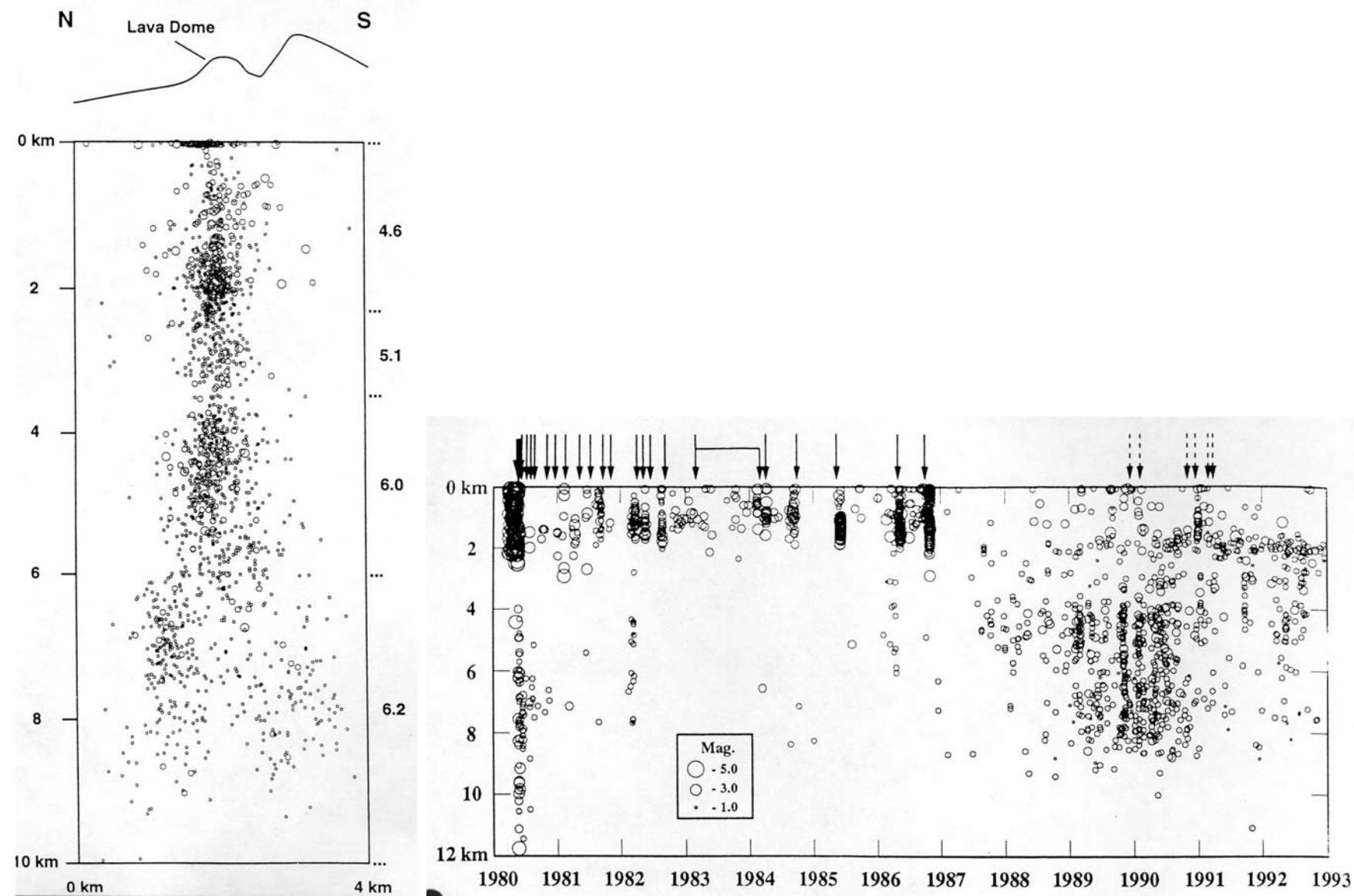
○ ≥ 10 km

} micro-earthquakes

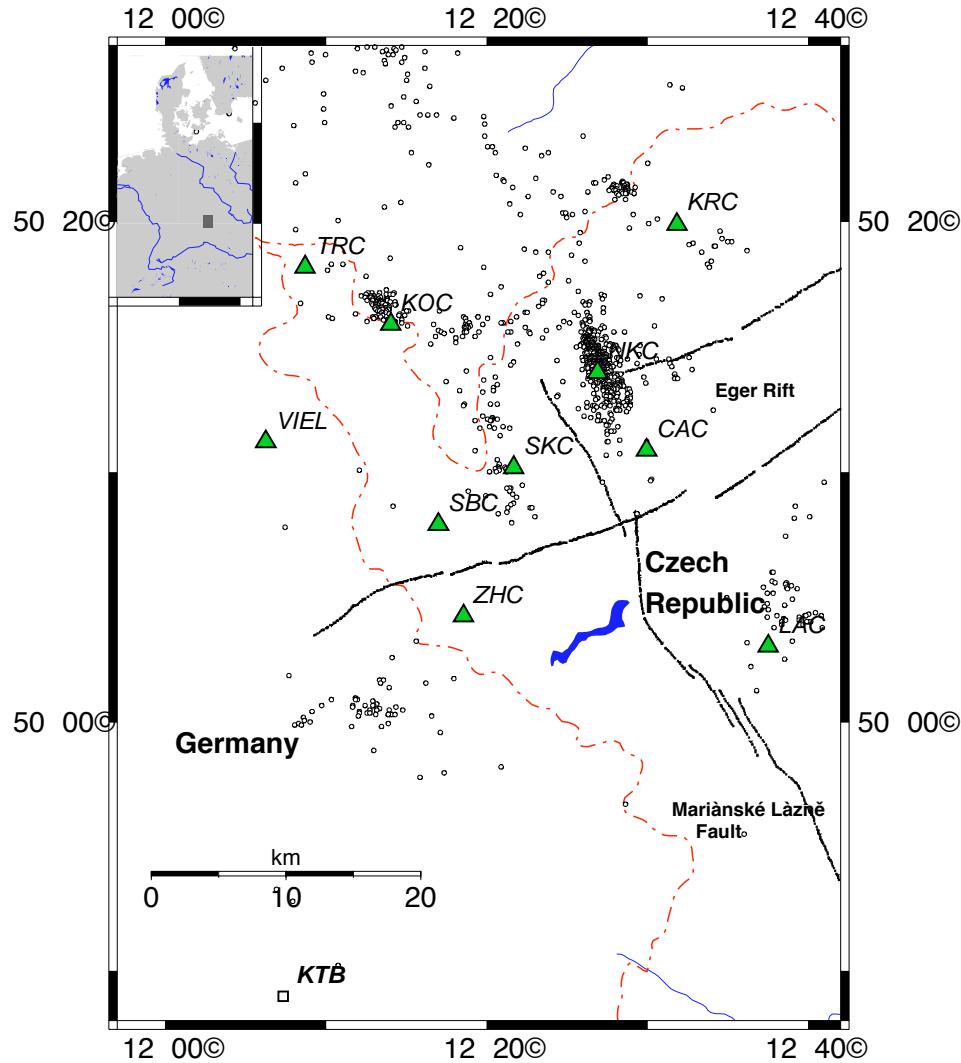
Eyjafjallajoekull 1994 dike intrusion



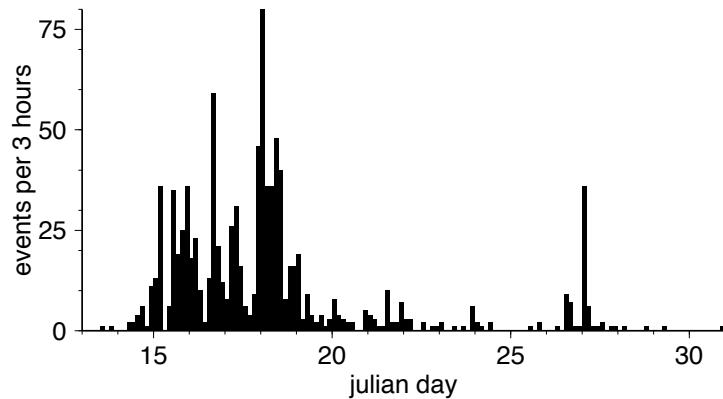
Mt St. Helens



The Vogtland swarm region



- ≈ 1800 events during two weeks
- $M_L \leq 3$
- tiny clustered source region

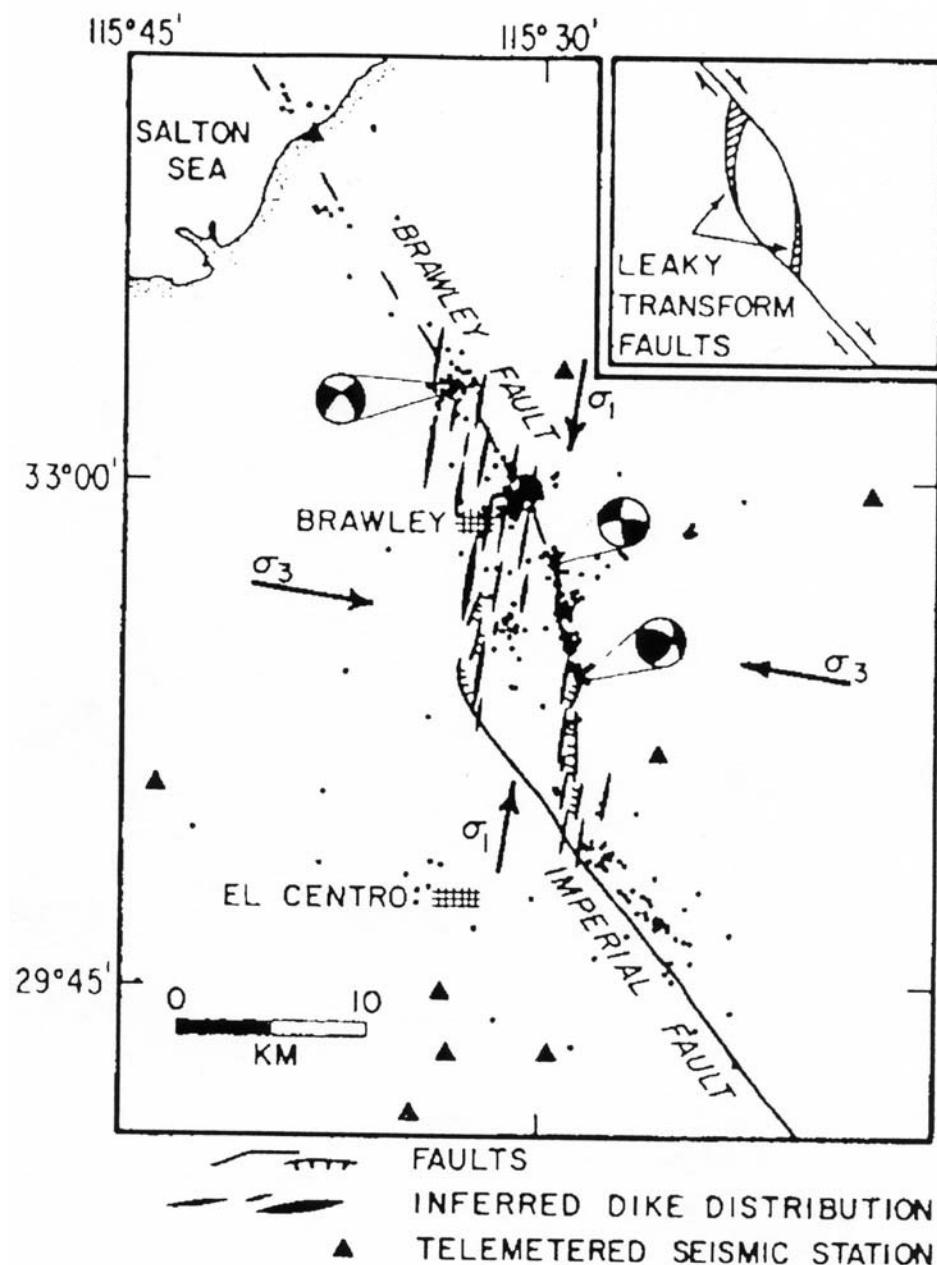


elements of swarm models

1. highly fractured media, without large, planar faults under shear
2. inhomogeneous stress fields and/or nonuniform loading
3. loading of cracks and faults by fluids

e.g. Mogi, 1963; Sammis & Julian, 1987; Rubin & Gillard, 1998

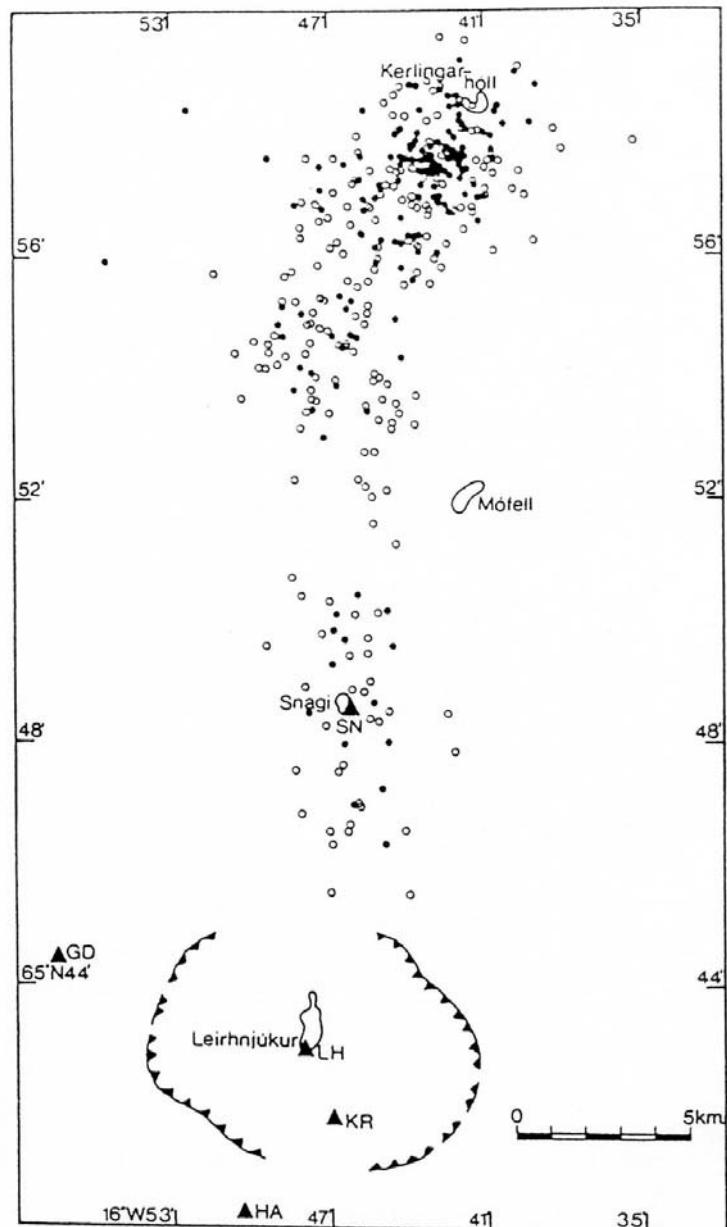
Hills model



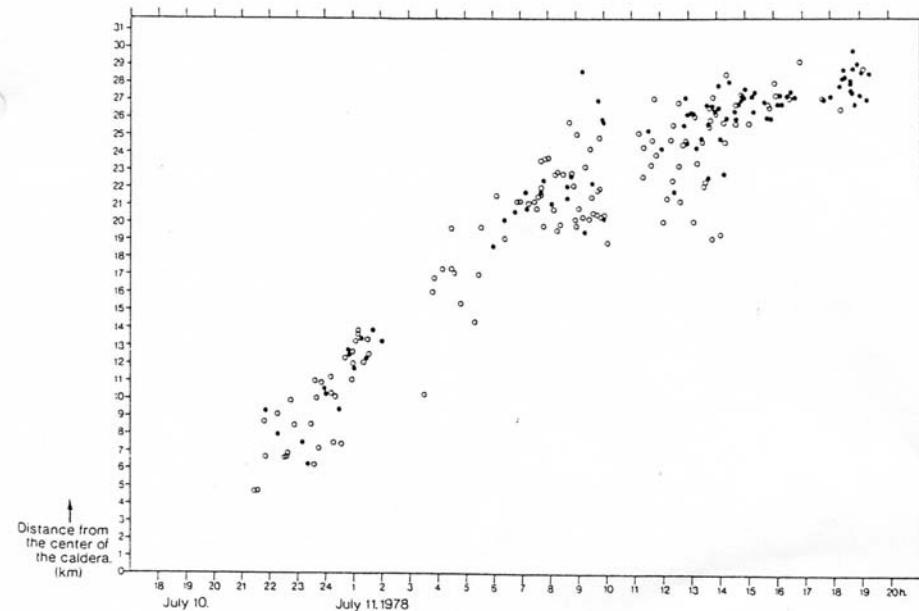
swarms are caused by stress loading (e.g. by dikes) in the extensional field of leaky transform faults.

Dikes and shear cracks use different faults.

Intrusion-induced swarms



Krafla 1978 swarm induced by lateral intrusion

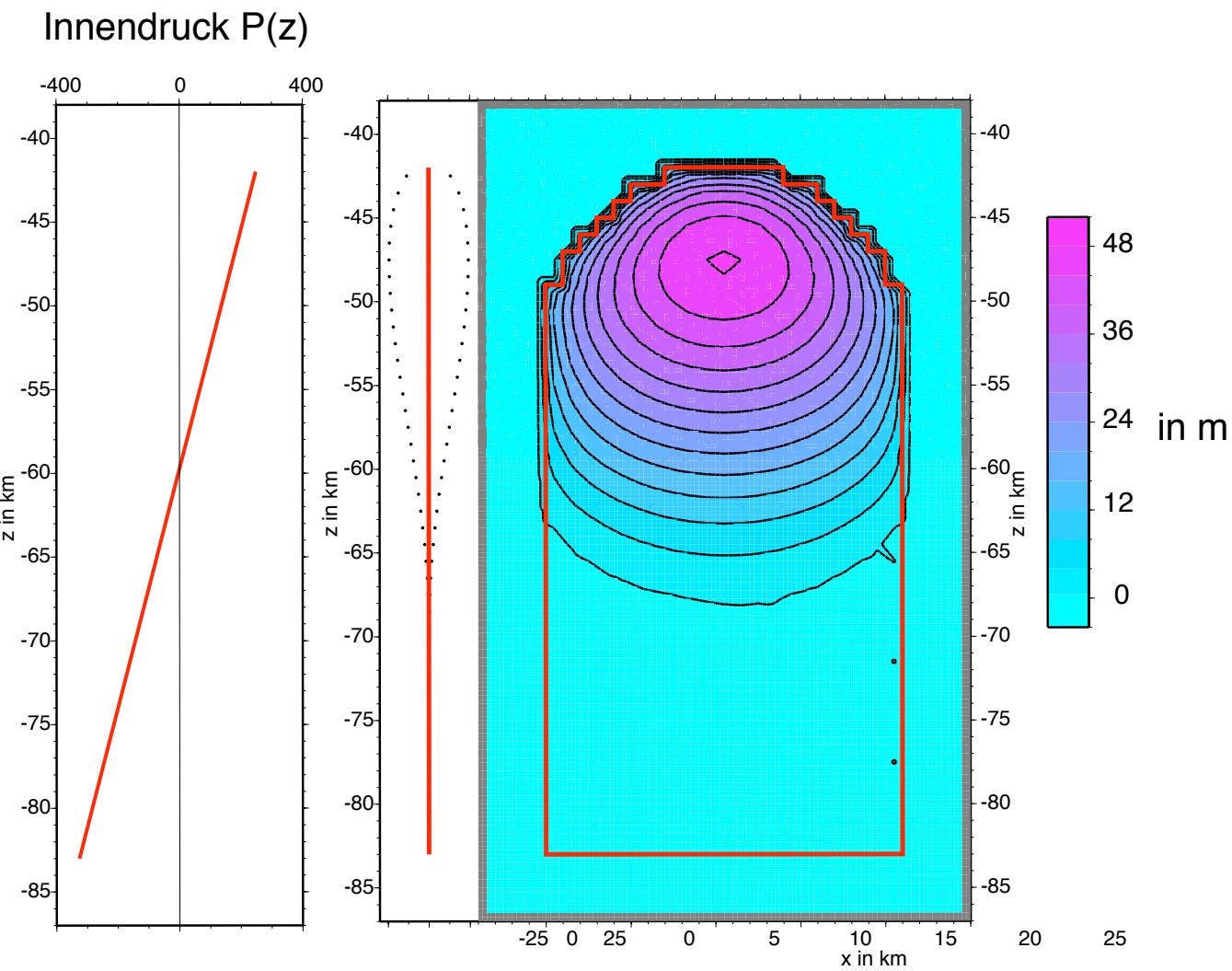


how do vertical intrusions look



numerical 3D buoyant dike

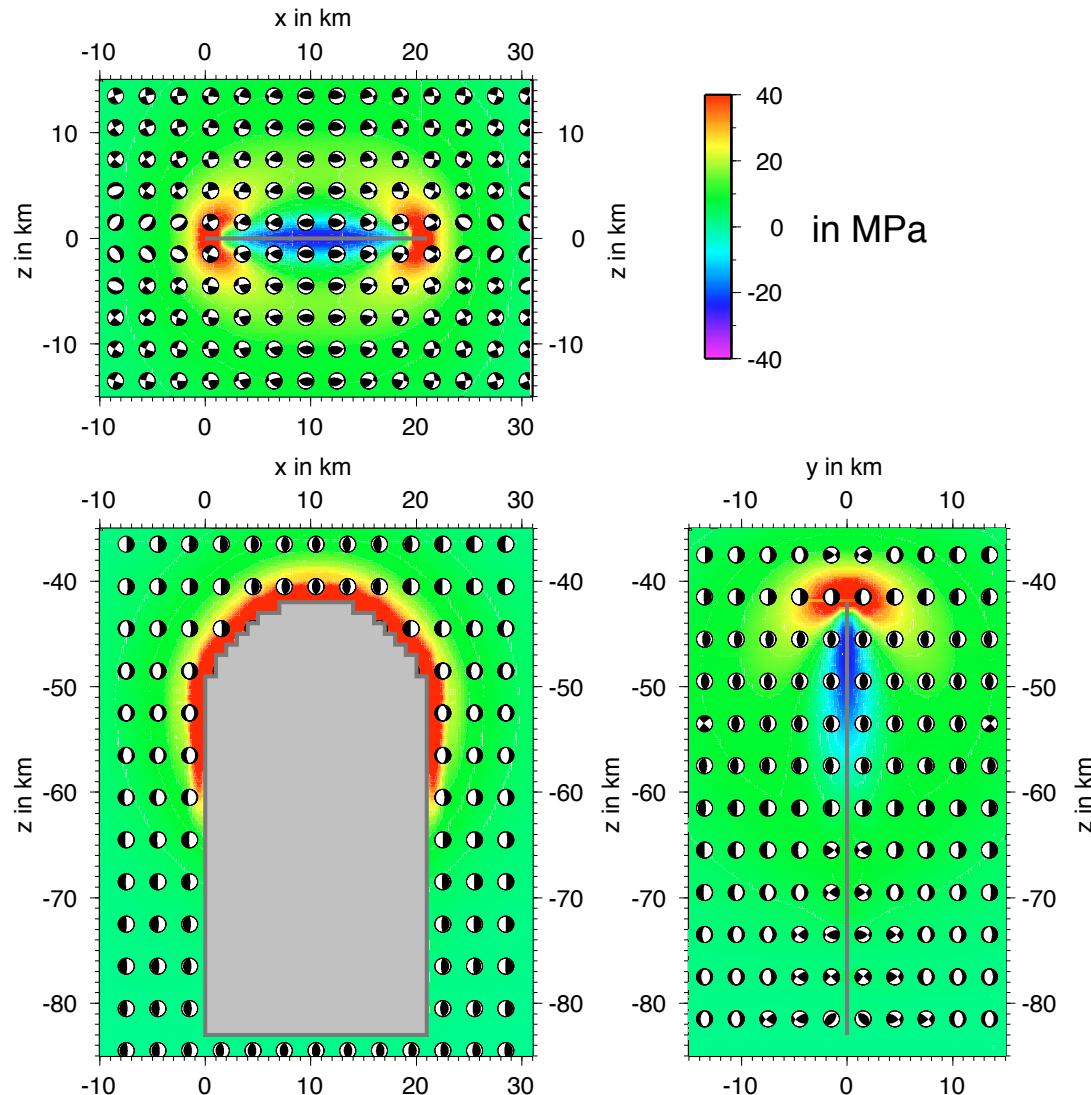
fluid-filled crack



stress loading by buoyant diking

Change of the Coulombstress

$$\Delta\sigma_C = \Delta\tau - \mu \Delta\sigma_n$$



Gelatine experiments

click

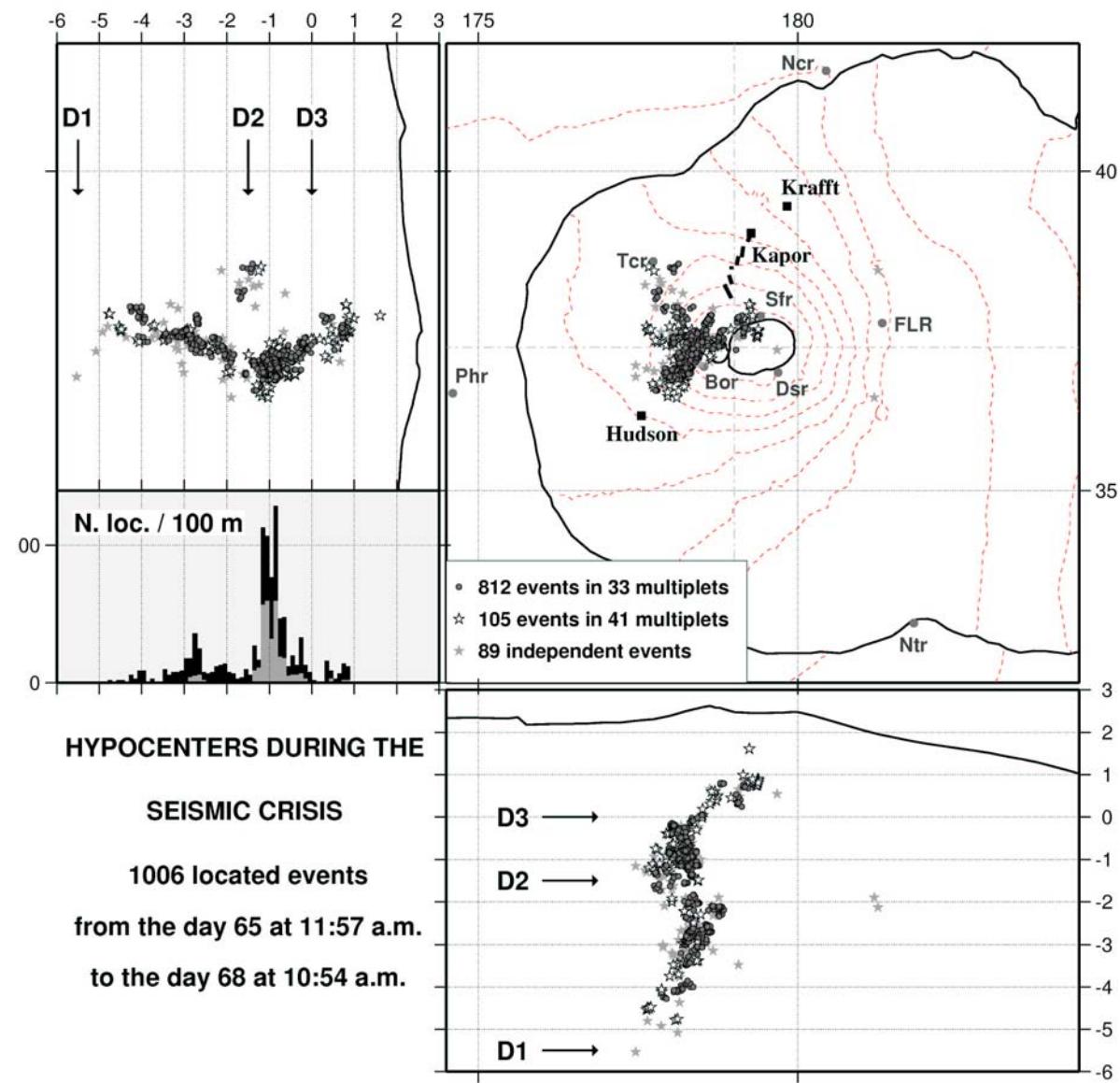


- 3% gelatine, brittle, transparent
- air is injected with a syringe
- characteristic length of cracks in the order of some cm



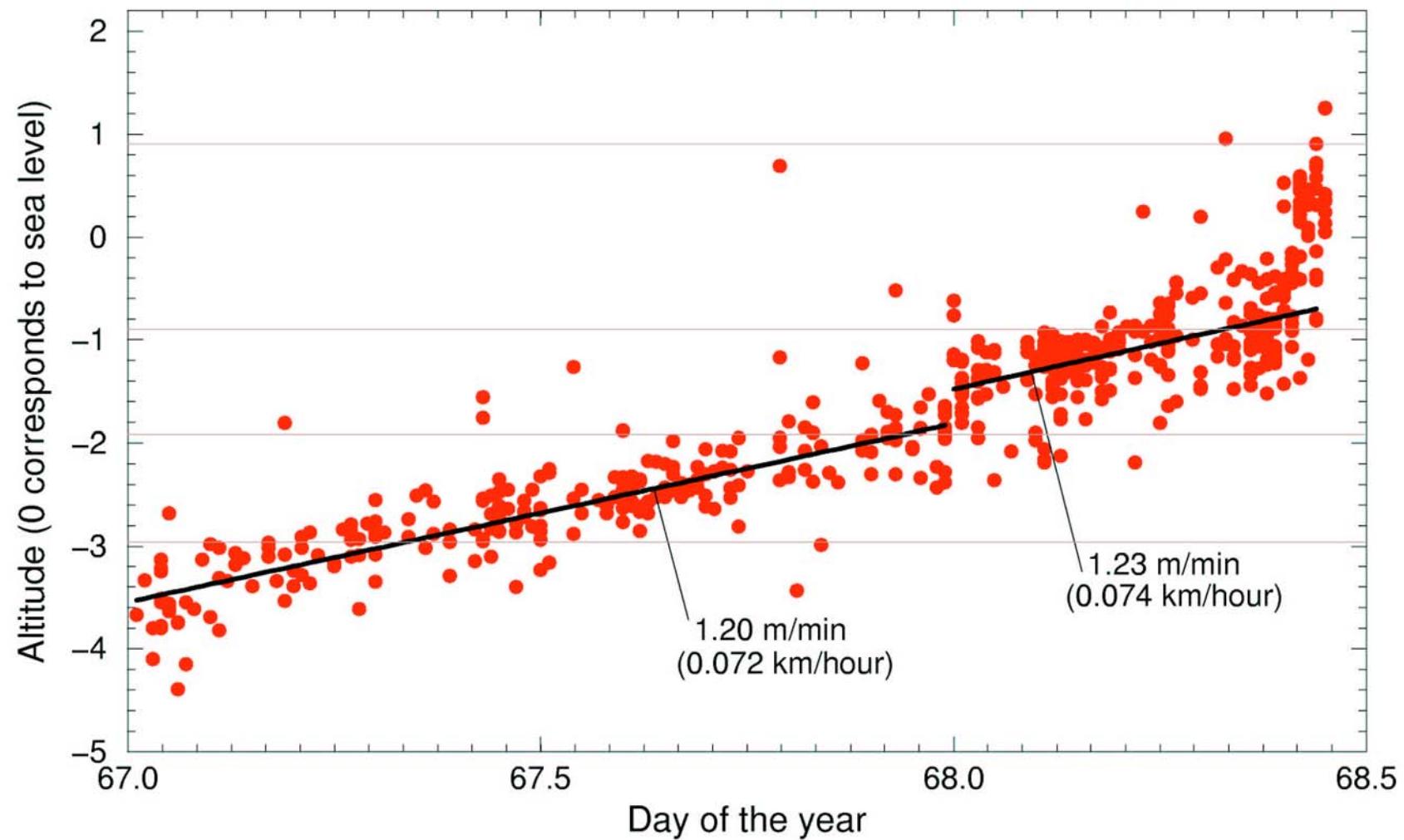
click speed of the movie: 25×

1998 Piton de la Fournaise eruption



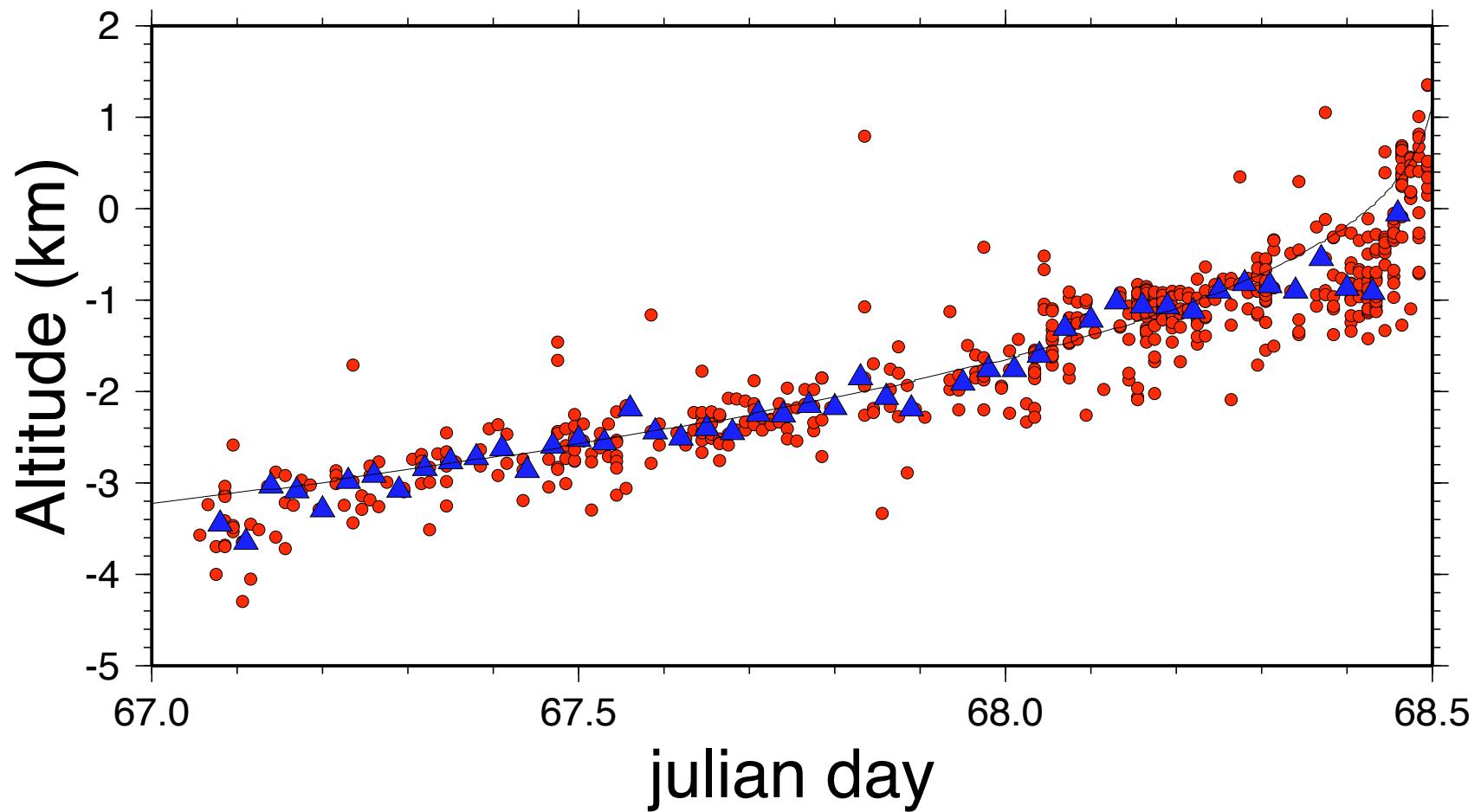
Battaglia (2001)

Migration of earthquake focal depths



from Battaglia (2001)

Crack tip predicted by gelatine model

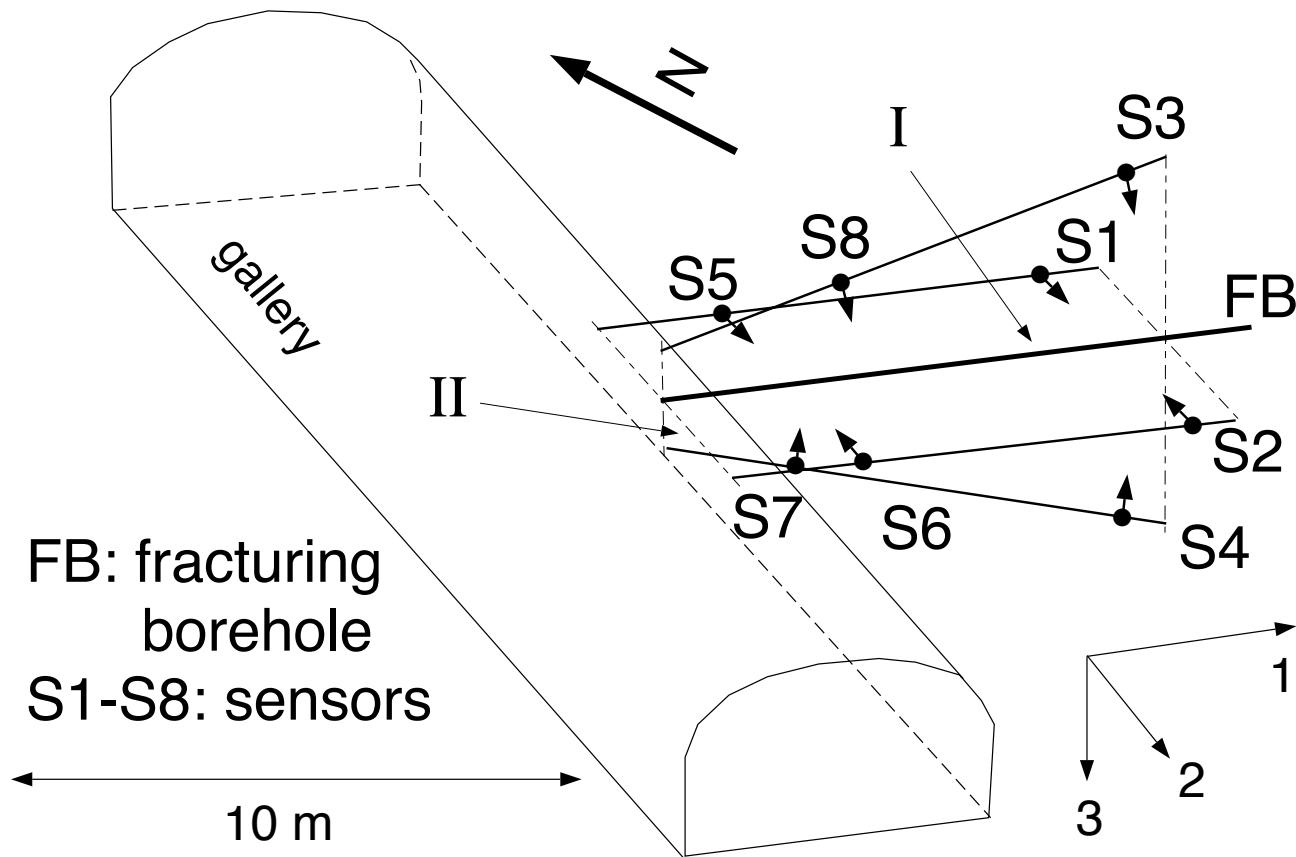


Rivalta & Dahm, submitted to GJI

non-buoyant intrusion

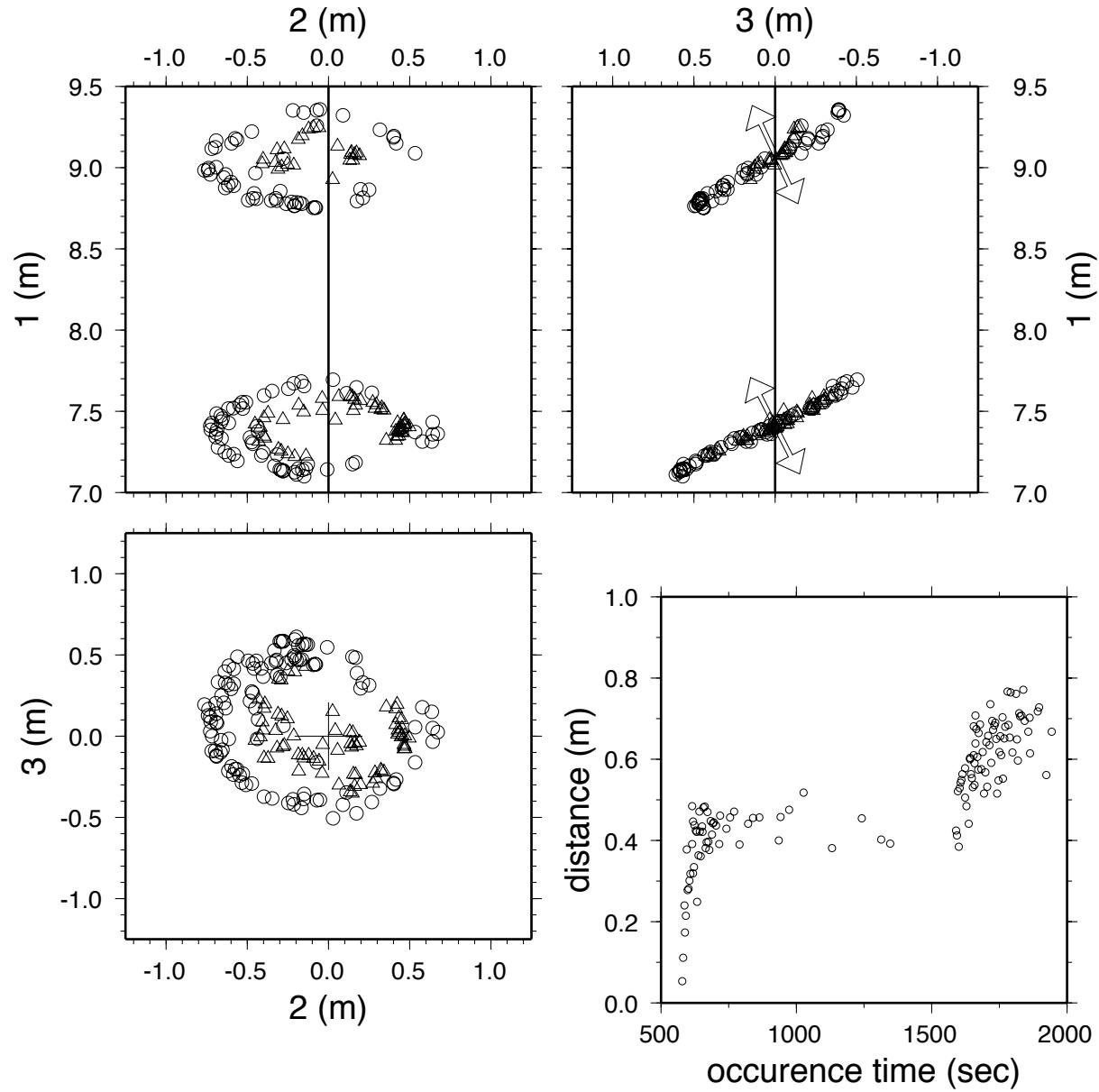
- Hydrofrac in Plexiglas: experiment by Prof. Rummel
- animated gif to follow incremental growth of sub-events

Example: Hydrofrac in a salt mine



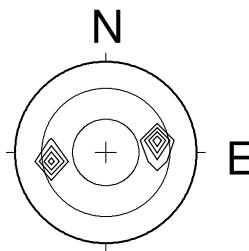
Dahm, Manthei & Eisenblätter (1999), Dahm (2001)

hypocenter follow penny shape crack

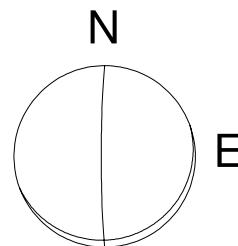


moment tensor solutions

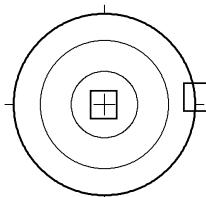
summed distribution
of P and T axis from
180 studied events



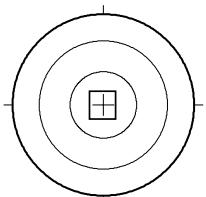
best double couple
of typical solution



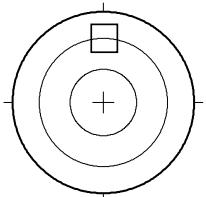
poles of typical nodal plane



poles of horizontal layers



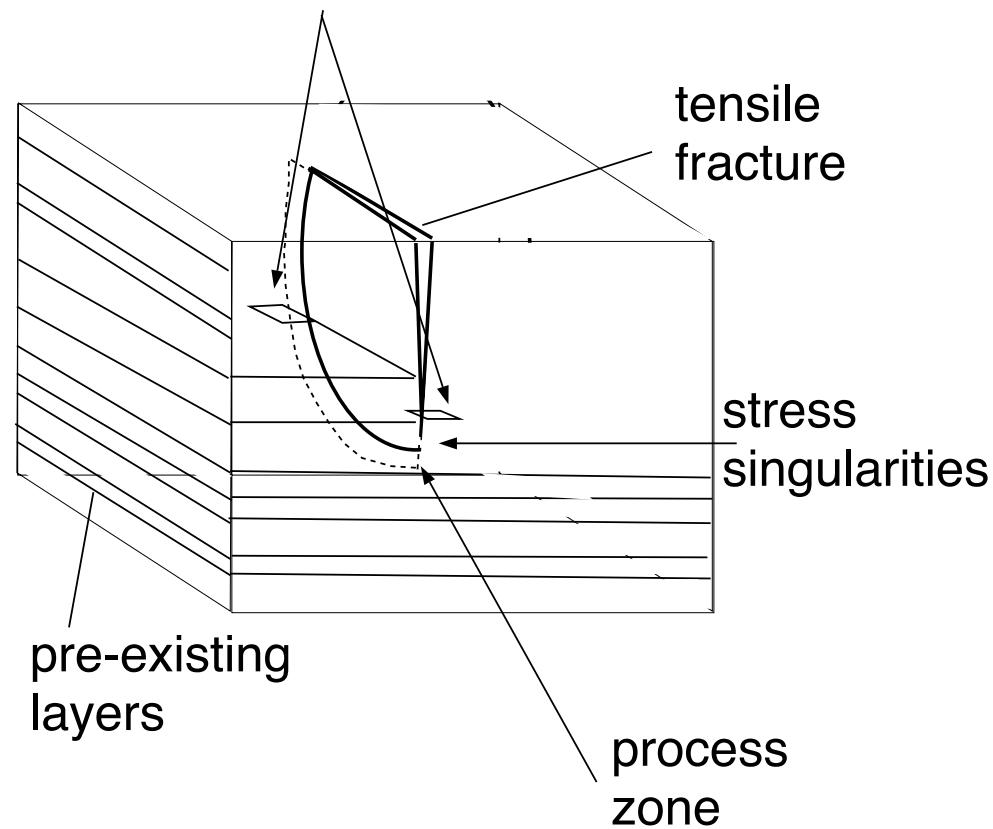
pole of fluid-filled fracture
plane



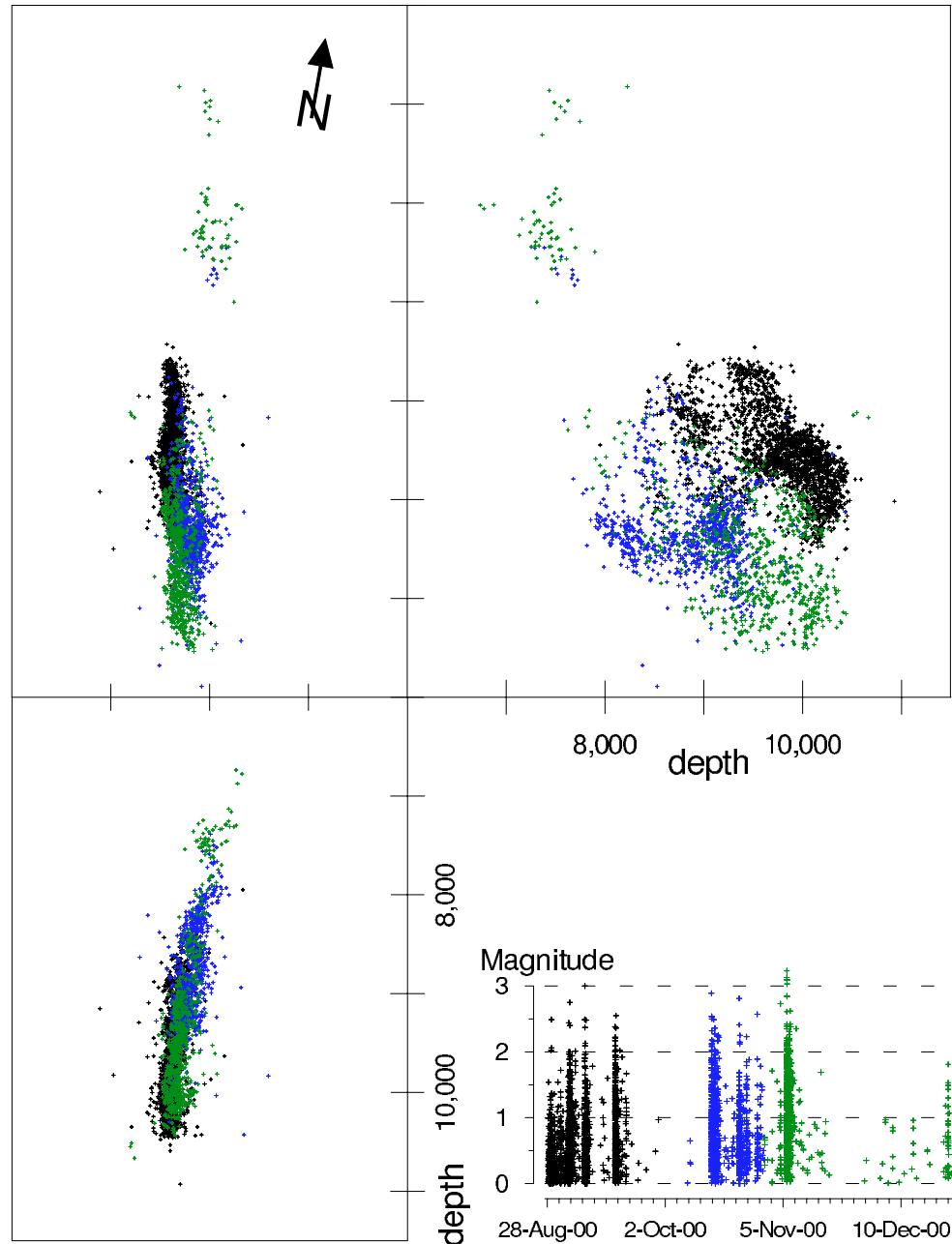
- shear cracks
- horizontal fault planes

Interpretation

micro cracks use horizontal planes of weakness



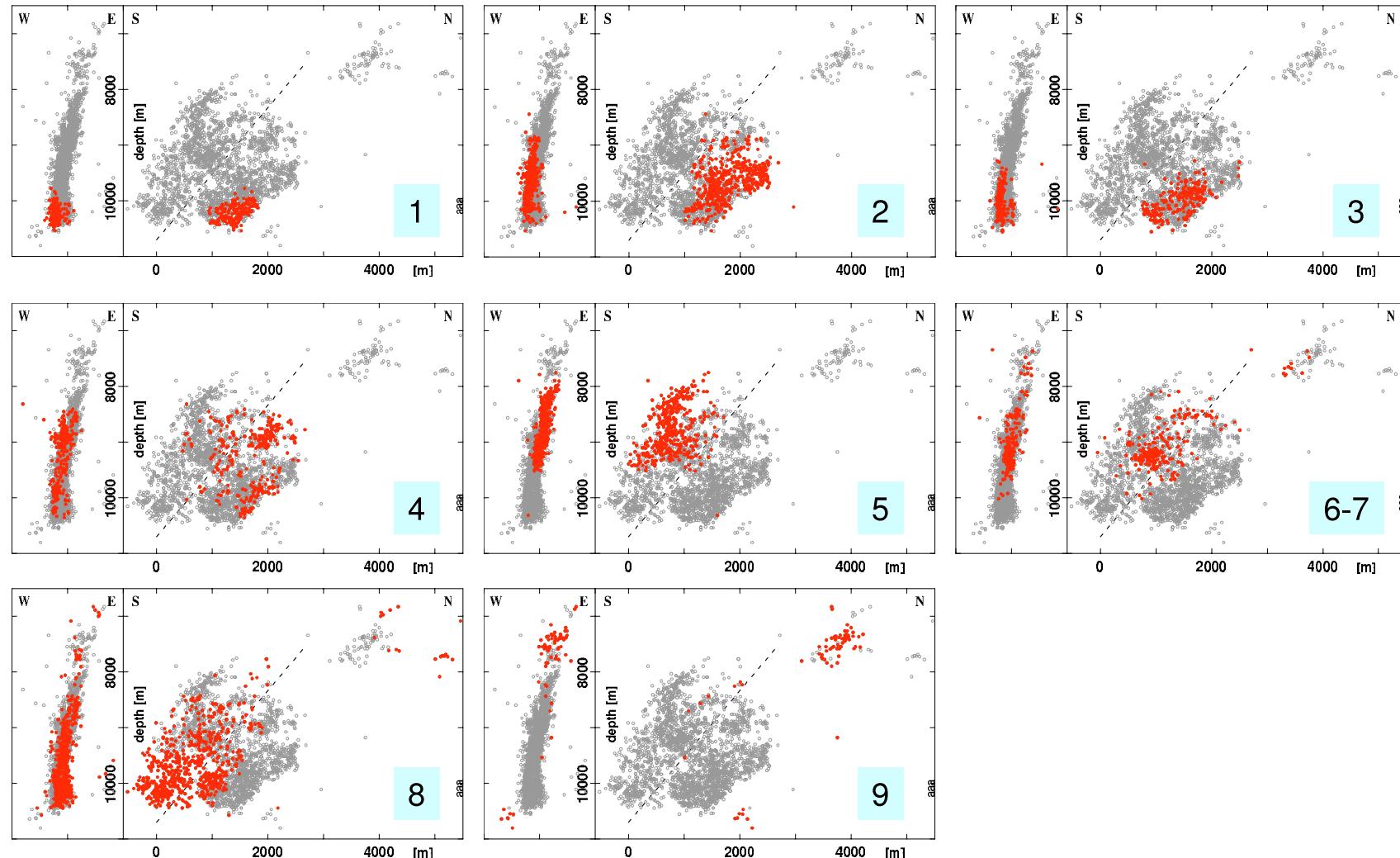
Example: Vogtland 2000 swarm



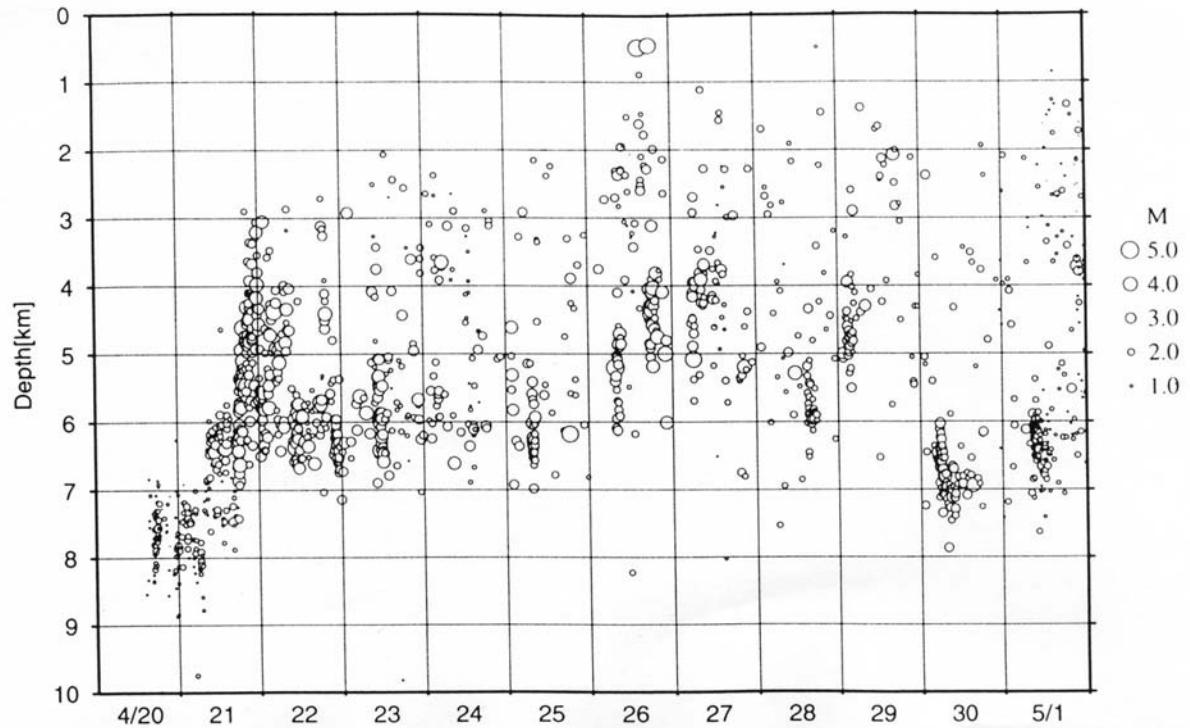
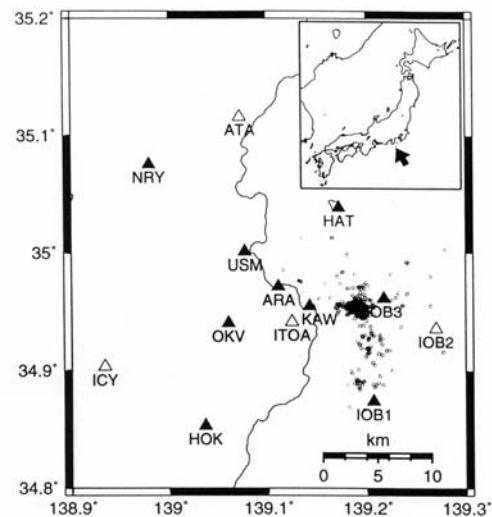
Fischer (2002)

Dahm, ICTP 2004 II – p.23/??

migration of events

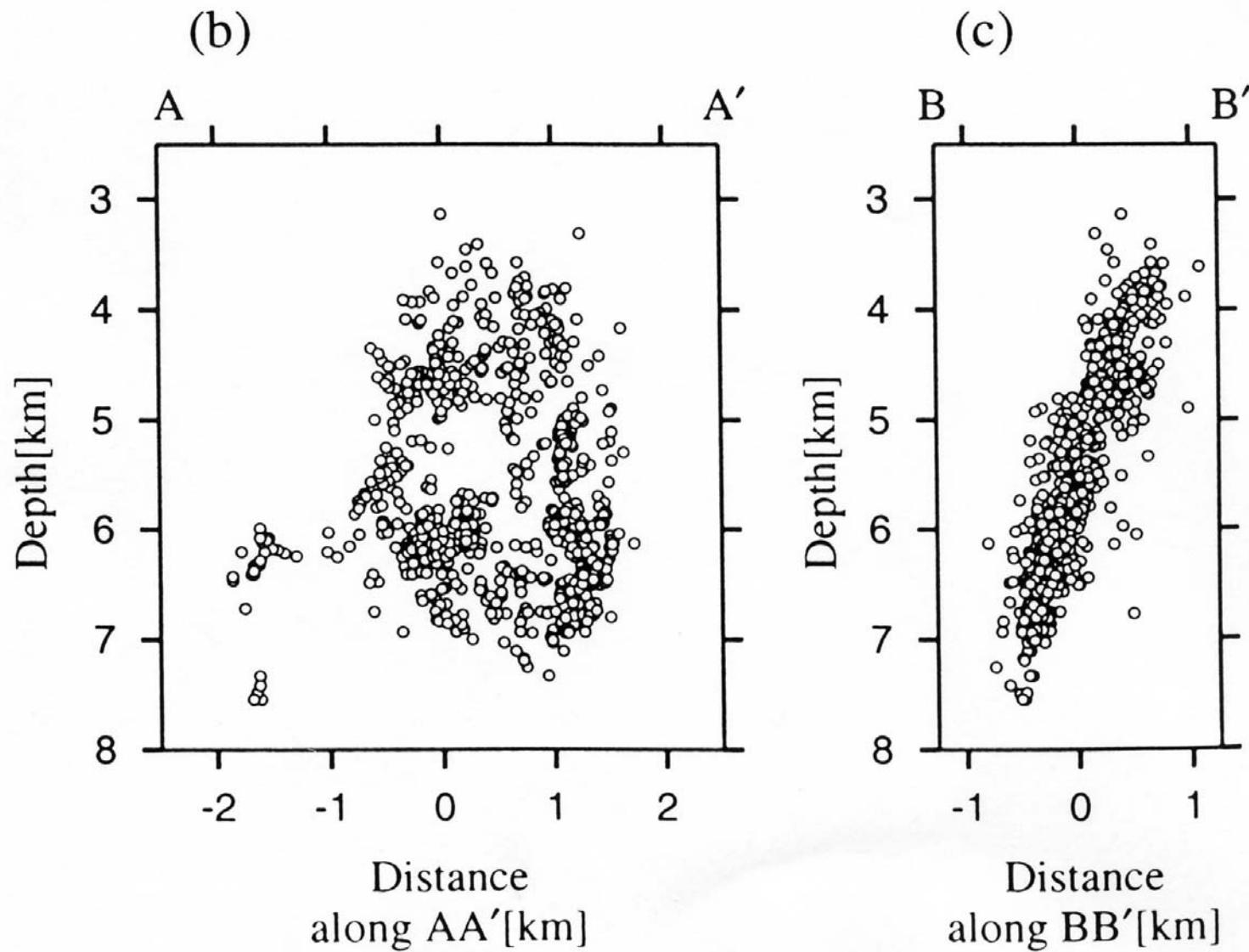


Example: Izu Bonin intrusion 2000



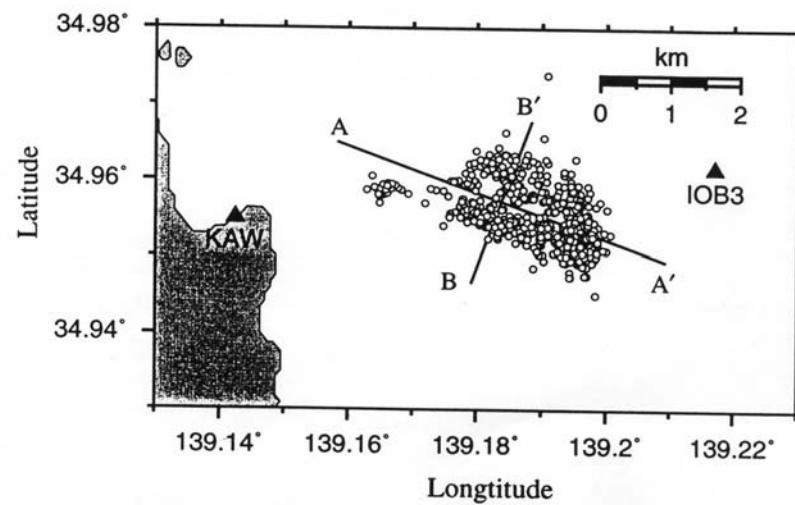
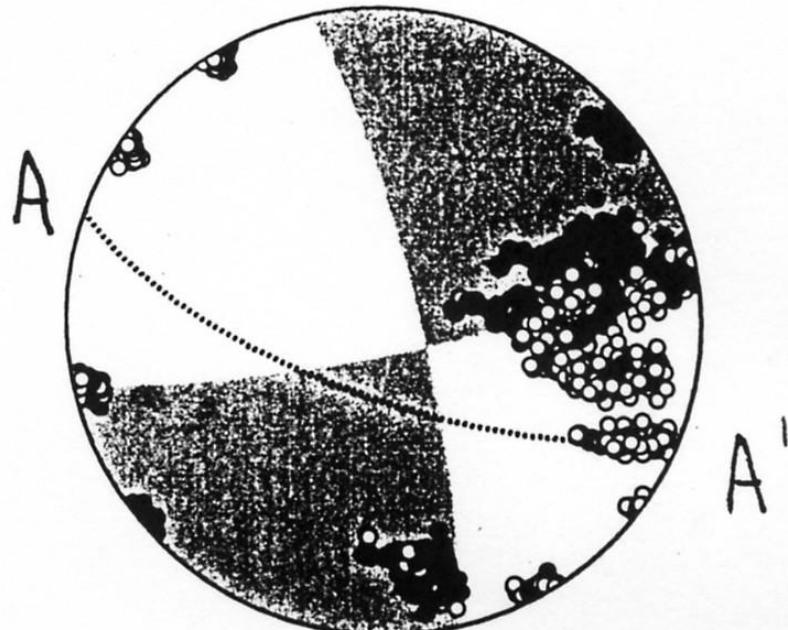
Hayashi & Morita, GJI, 2002: A magma intrusion process inferred from hypocenter migration of earthquake swarm

Izu Bonin



Source mechanism study

(b)



Composite focal solution of similar events indicate that shear cracks are induced outside the dike volume

Summary

1. high-precision relative location shows migration of events and can reveal intrusion history
2. (relative) moment tensor are important to understand intrusion physics
3. intrusion-induced earthquakes are usually (off-plane) shear events. They may have small volume expansion components