



the
abdus salam
international centre for theoretical physics

40¹⁹⁶⁴
anniversary
2004

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**"7th Workshop on Three-Dimensional Modelling
of Seismic Waves Generation and their Propagation"**

25 October - 5 November 2004

**Theoretical and Observed Envelopes
of Scattered High-Frequency Seismic Waves
at Local to Regional Distances**

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A.A. Gusev

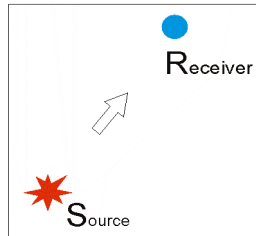
**Theoretical and observed envelopes
of scattered high-frequency seismic
waves at local to regional distances**

OUTLINE:

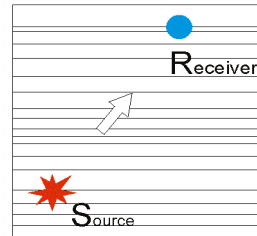
1. RANDOM MEDIA, RANDOM AND OBSERVED SIGNAL
2. MORPHOLOGY OF SCATTERED WAVES ON THE EARTH. CODA
3. THEORY. RANDOM SCATTERERS, RANDOM INHOMOGENEITY
4. SIMULATED ENVELOPES
5. INVERSION FOR TURBIDITY
6. NON-UNIFORMITY OF SCATTERER DENSITY IN THE EARTH

Common models of the medium where the waves propagate

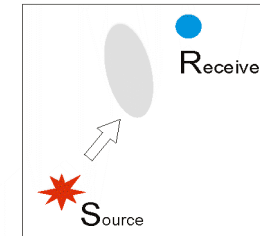
DETERMINISTIC MEDIA



UNIFORM

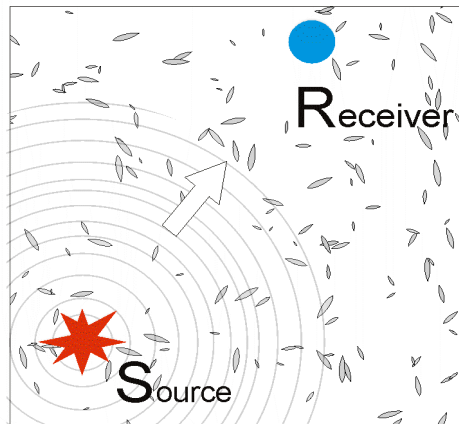


LAYERED HALF-SPACE

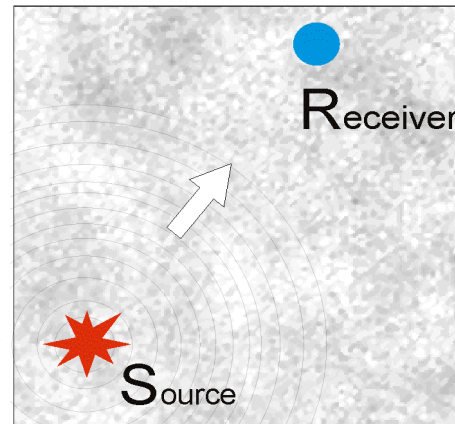


DETERMINISTIC OBSTACLE

RANDOM MEDIA



RANDOM DISTRIBUTION
OF OBSTACLES/SCATTERERS



RANDOM FIELD
OF PROPERTIES
(λ, μ, ρ)

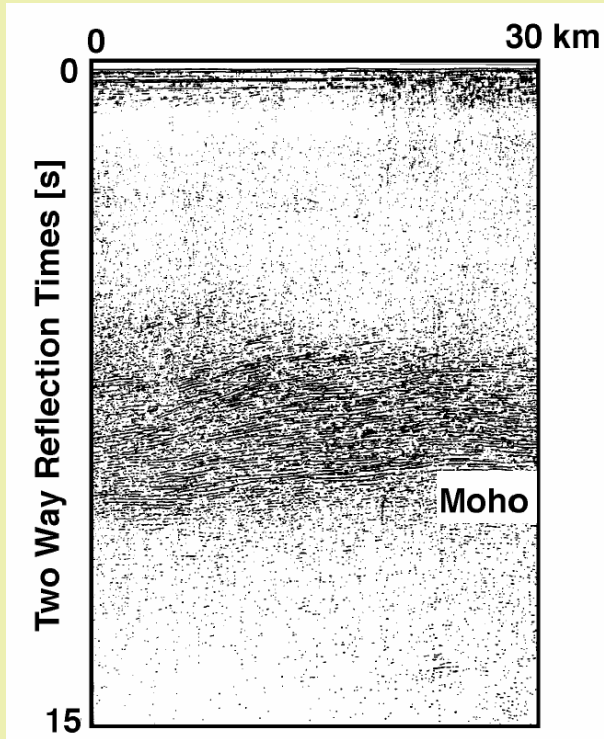
RANDOM PERTURBATION OF PROPERTIES

$$\begin{aligned}\lambda(\mathbf{x}) &= \lambda_0(1 + \varepsilon_\lambda(\mathbf{x})); \\ \mu(\mathbf{x}) &= \mu_0(1 + \varepsilon_\mu(\mathbf{x})); \\ \rho(\mathbf{x}) &= \rho_0(1 + \varepsilon_\rho(\mathbf{x}))\end{aligned}$$

Weak inhomogeneity
 $\varepsilon \ll 1$

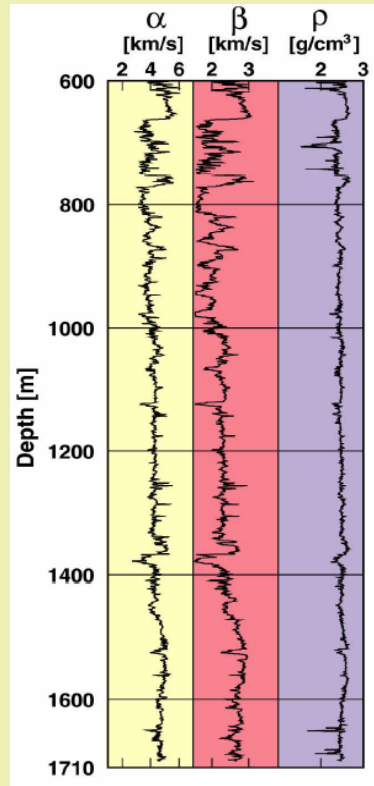
Acoustic case
 $c(\mathbf{x}) = c_0(1 + \varepsilon(\mathbf{x}))$
Coefficient of refraction
 $n(\mathbf{x}) = (1 + \varepsilon(\mathbf{x}))$

Random-like real-Earth structures



Example reflection-seismic section: strong heterogeneity in the lower crust (Warner, 1990)

anisotropic
non-uniform
random field



**Example well log
Persistent oscillation
of elastic parameters
(Shiomi et al.,1997)**

non-Gaussian
random field

RANDOM INHOMOGENEITY OR
PERTURBATION OF PROPERTIES:

$$\lambda(\mathbf{x}) = \lambda_0(1 + \varepsilon_\lambda(\mathbf{x}));$$

$$\mu(\mathbf{x}) = \mu_0(1 + \varepsilon_\mu(\mathbf{x}));$$

$$\rho(\mathbf{x}) = \rho_0(1 + \varepsilon_\rho(\mathbf{x}));$$

Background: λ_0, μ_0, ρ_0

Perturbation: $\varepsilon_\lambda(\mathbf{x}), \varepsilon_\mu(\mathbf{x}), \varepsilon_\rho(\mathbf{x})$

Acoustic case: $c(\mathbf{x}) = c_0(1 + \varepsilon(\mathbf{x}));$

Usual assumptions w.r.t. perturbation field:

(1) Weak: $\varepsilon(\mathbf{x}) \ll 1$

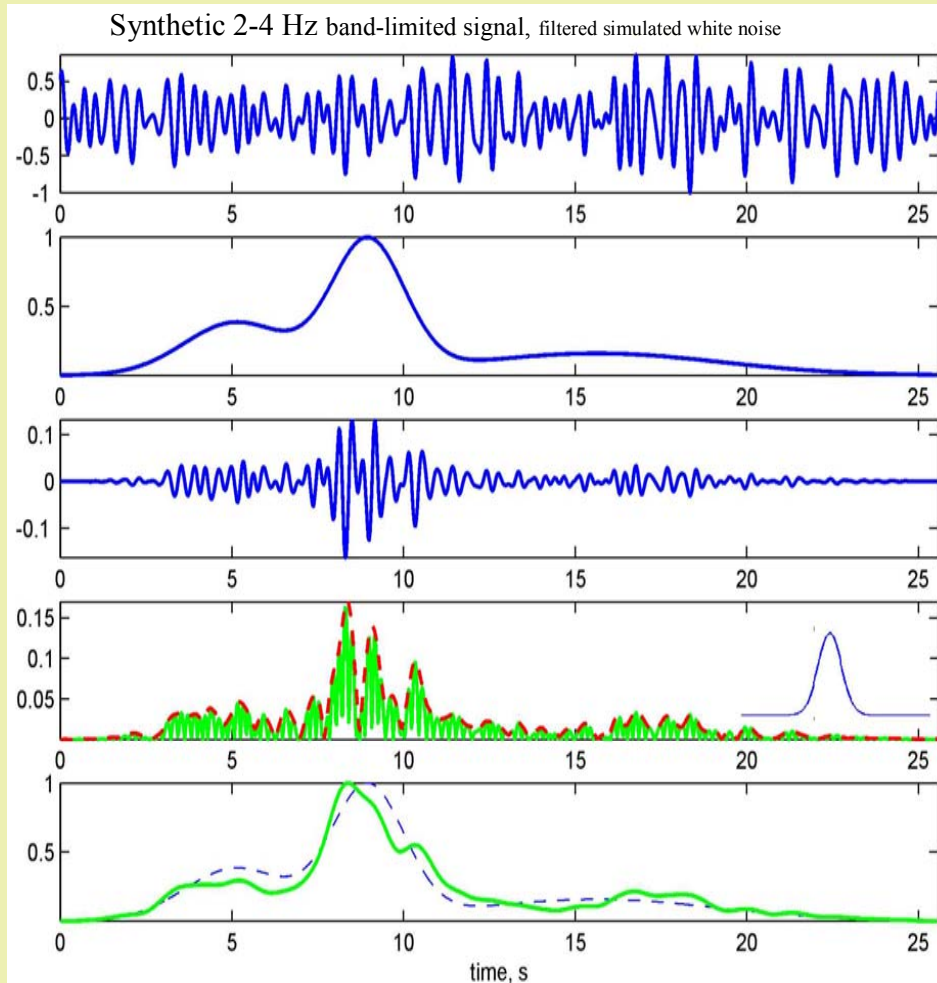
(2) Uniform = homogeneous = stationary:

$$\text{Cov}(\varepsilon(\mathbf{x}), \varepsilon(\mathbf{x} + \mathbf{y})) = \sigma_\varepsilon^2 \rho(\mathbf{y})$$

(3) Isotropic: $\rho(\mathbf{y}) \rightarrow \rho(\|\mathbf{y}\|) = \rho(r)$

(in the non-Gaussian case,
more details are needed)

Random signal, envelope, power (1)



— stationary random signal $x(t)$
 constant mean power or variance: $\sigma^2(t) = \langle x^2(t) \rangle$
 constant “true” rms amplitude: $a_{rms}(t) \equiv \sigma(t)$

— “true” envelope or
 modulating function $a(t)$
 $(a^2(t) - \text{“True” power time history})$

— $y(t) = x(t) \times a(t)$: simulates
 observed QUASISTATIONARY signal,

— $\text{abs}(y(t))$
 - - - - - module of analytic signal (MAS)

— $a_e(t) = \text{SQRT}(\text{smoothed } y^2(t))$
 - - - - - $a(t)$

$a_e(t)$: empirical envelope,
 an estimate for “true” $a(t)$,
 like those derived from data

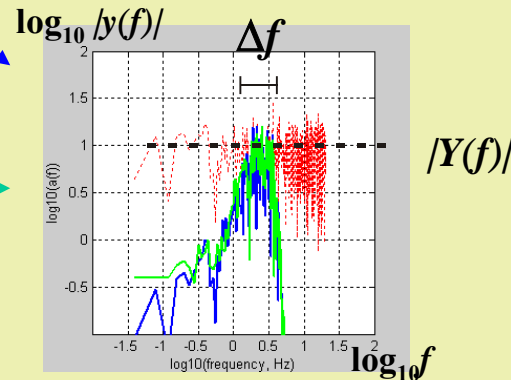
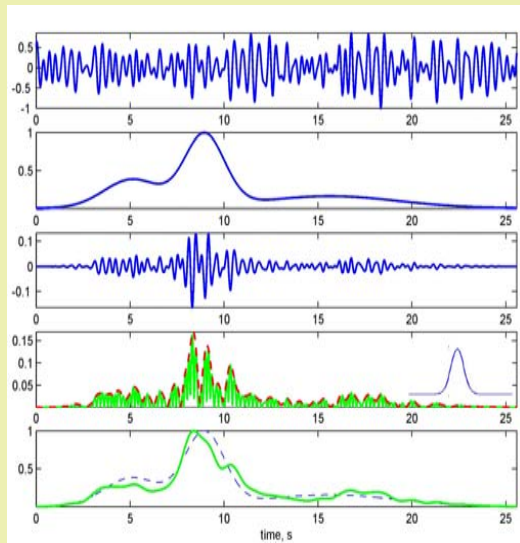
$a_e^2(t)$: observed time history
 for power

$a_e(t)$ can be also estimated using signal peaks

“True”: pertains to ENSEMBLE AVERAGE or MEAN of the process

“Observed”: pertains to a single SAMPLE FUNCTION or
 a REALIZATION of the random process

Random signal, envelope, power (2)



3. Denote $P(f/t)$ signal power spectrum, average over a window of length d around t

Then

$$P(f/t) = 2|y(f)|^2 / d$$

1. Main signal parameters:

f_c – central frequency of a band

Δf – bandwidth ($1/\Delta f$ - time scale of “instant” power change)

$t_{drift} \approx \max(a(t))(da(t)/dt)^{-1}$
– time scale of non-stationarity

T_{sm} – width of smoothing window

Condition of quasi-stationarity:

$$t_{drift} \gg 1/\Delta f$$

Condition on smoothing window:

$$T_{sm} \gg 1/\Delta f$$

2. Denote:

$|Y(f)|$ – Fourier amplitude spectral level, average over the bandwidth Δf

d – signal duration (or window duration)

y_{rms} – rms signal amplitude over d

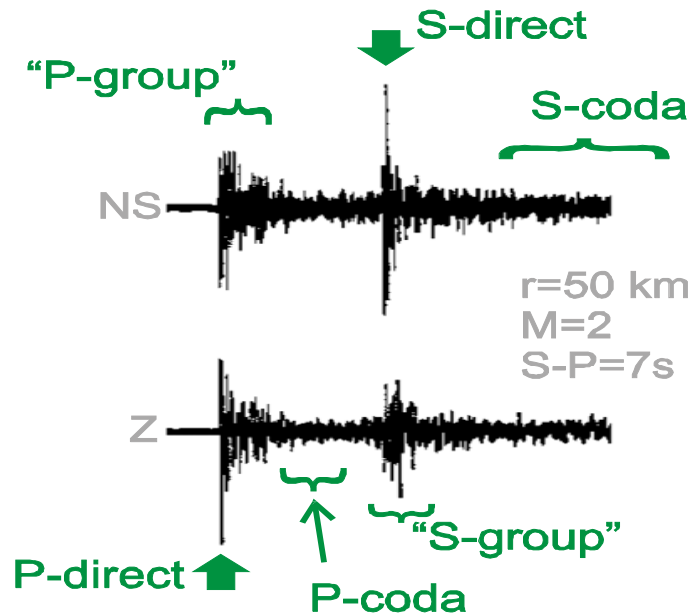
Then (Parseval’s theorem):

$$2 |Y(f)|^2 \Delta f = y_{rms}^2 d$$

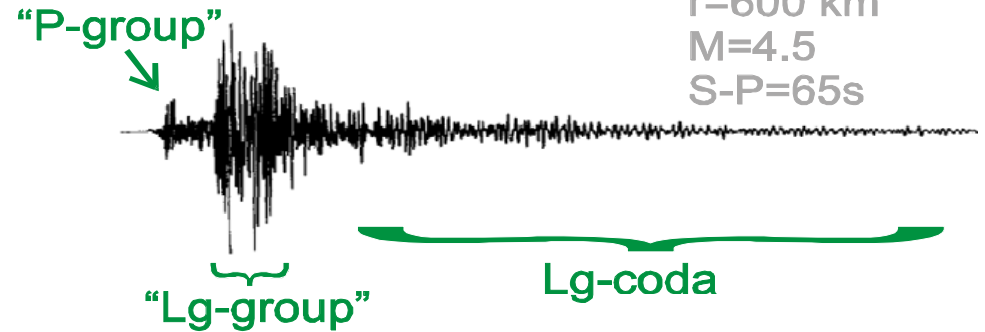
[permits to convert time domain to spectral domain estimates and back]

Regional seismograms – examples, morphology

local event



regional event



Maeda&Walter 1996

P-direct, S-direct – represent source-time-function, disappear at $r=15-70$ km for shallow events, short spikes for low magnitudes

P-group – appearance defined by medium, mix of P-direct, P-P forward scattered and P-S converted

P-coda – P-P wide-angle scattered and P-S converted

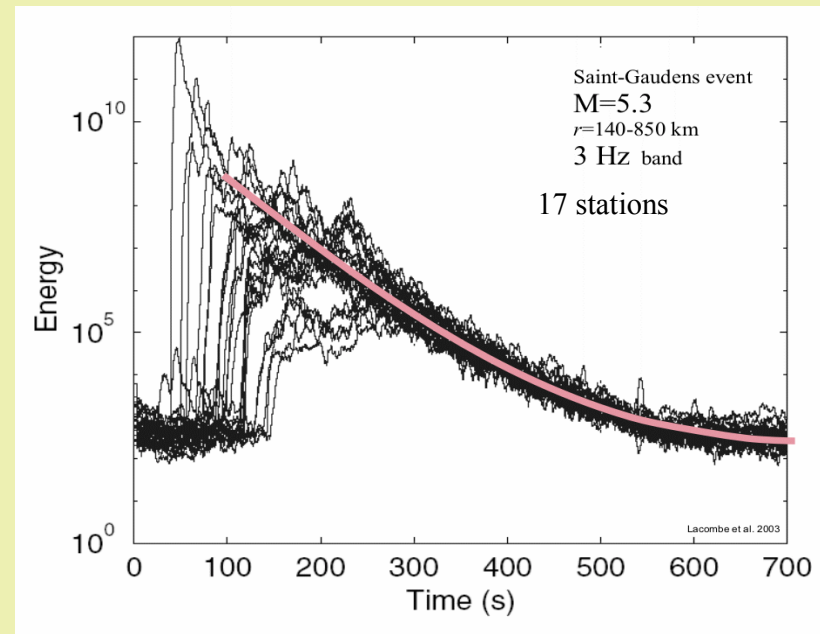
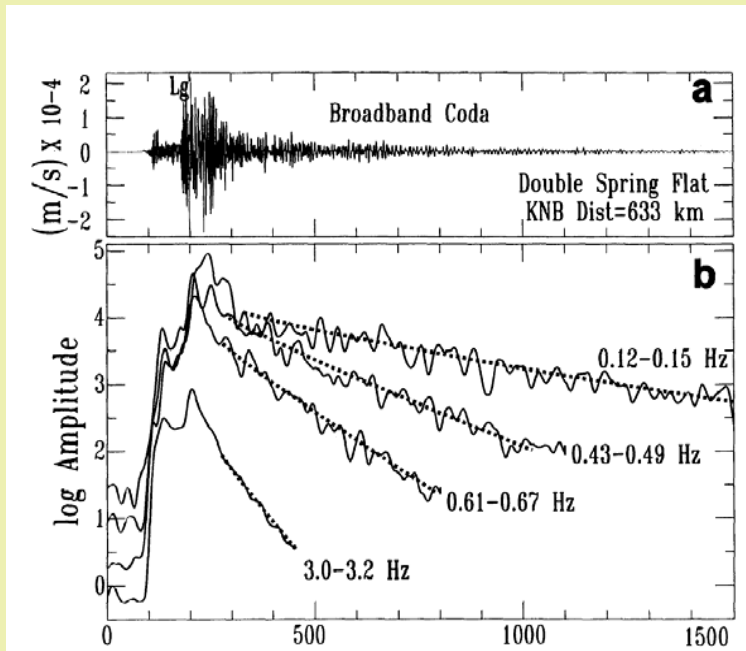
S/Lg-group – mix of S and HF surface waves, direct and forward-scattered

S/Lg-coda – S and HF surface waves, wide-angle scattered

r=180 km
M=3.5
S-P=20s

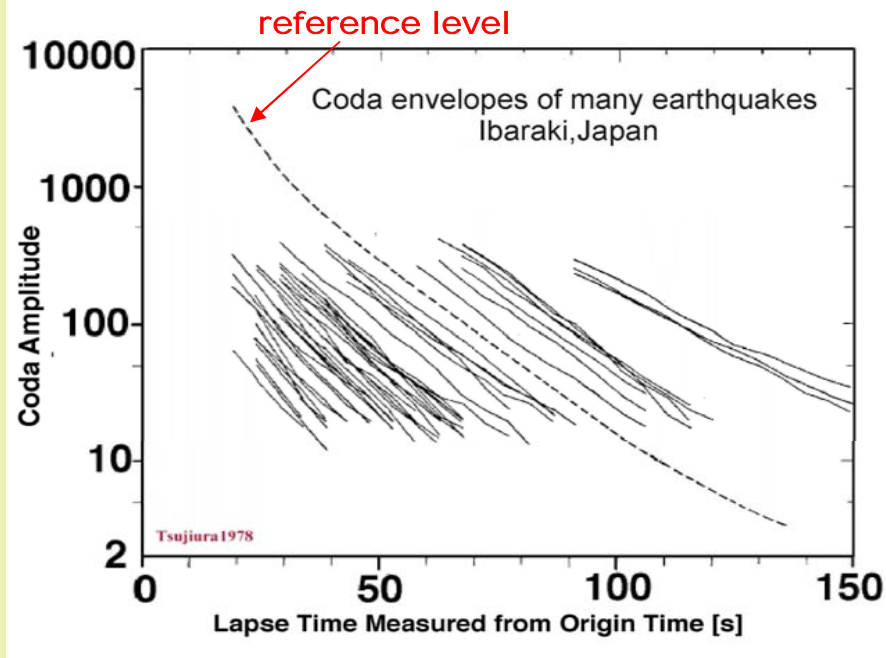


Regional envelopes

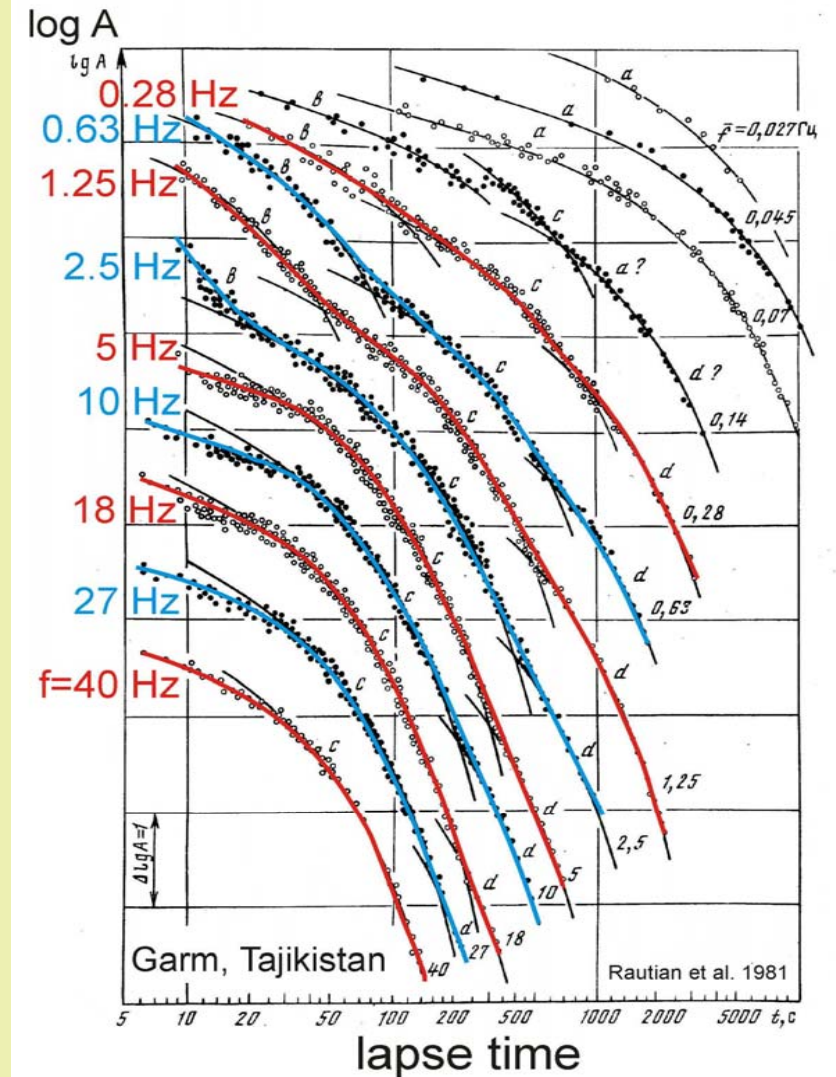


1. Envelopes from band-filtered HF records show systematic structure, first of all coda
2. To select coda, use sufficient delay, like $2t_s$ (*coda window*)
3. Coda decay is monotonous, regular, frequency dependent
4. Coda envelope is approximately station-independent
(a certain constant factor is present, it depends on local geology, useful for site specification)

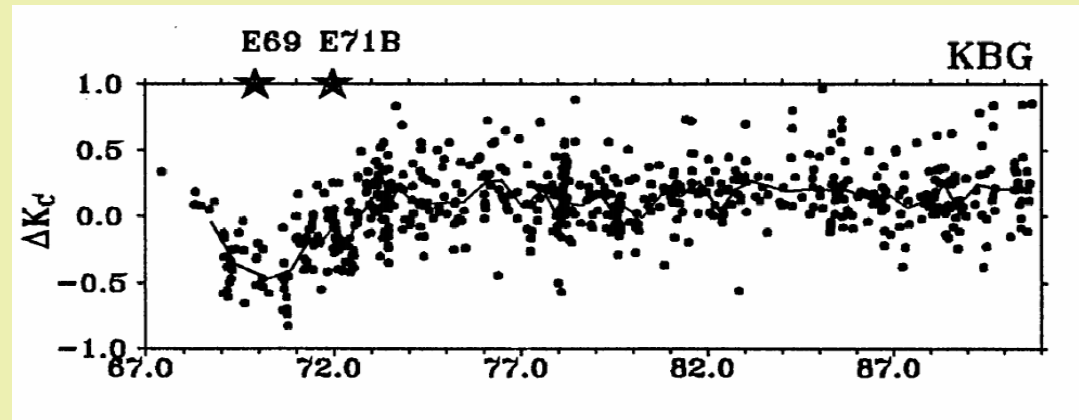
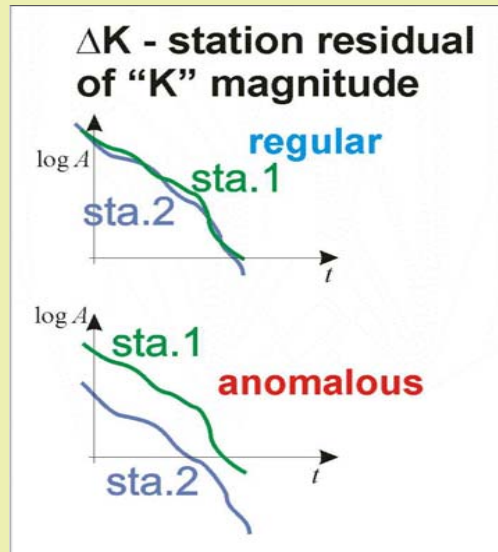
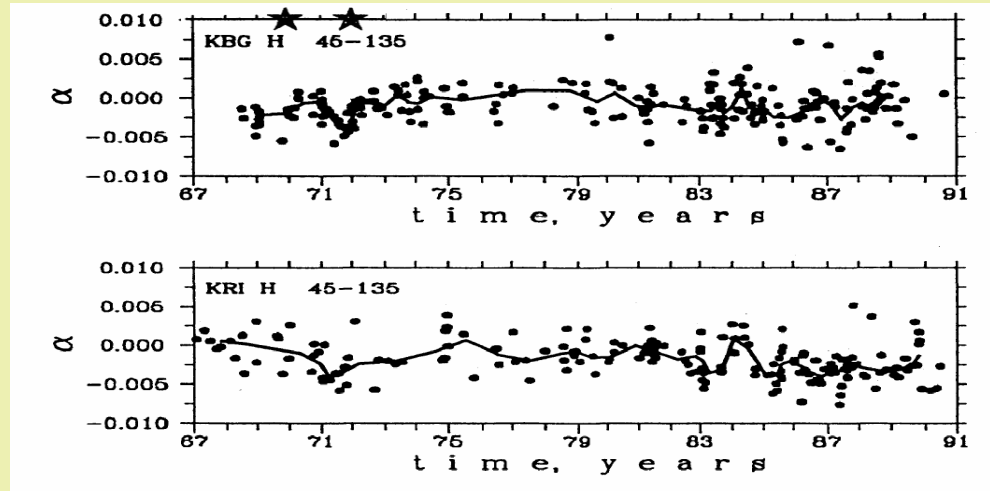
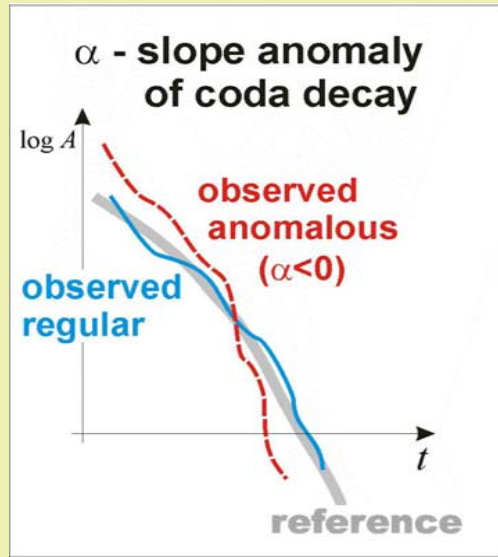
Regional coda



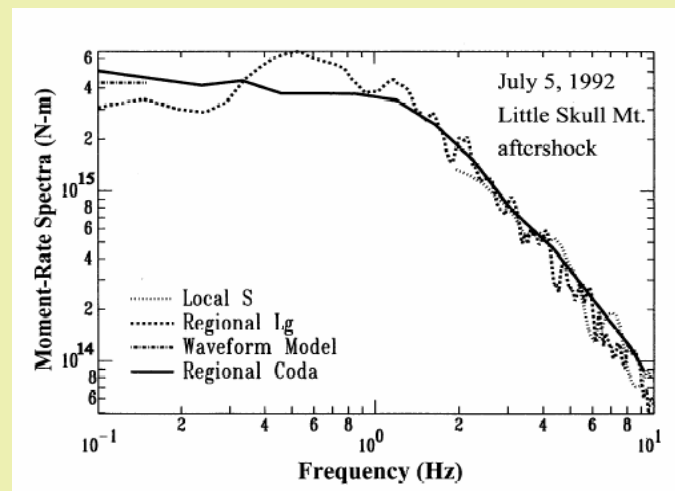
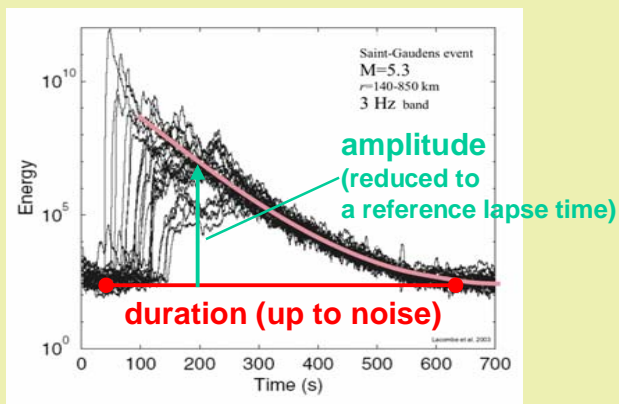
1. Coda envelope *shape* is approximately event-independent.
2. The scaling factor to reduce observed coda amplitude to a reference level gives (f -dependent) coda magnitude. After additional calibration it gives source spectrum $\dot{M}_0(f)$



Temporal variations of coda shape and level

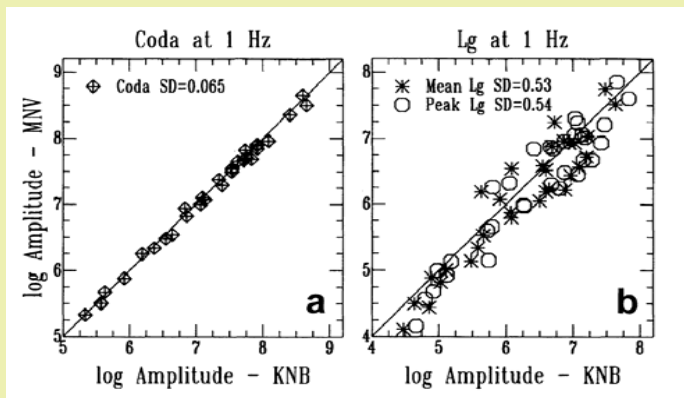


Coda magnitudes. Source spectra from coda



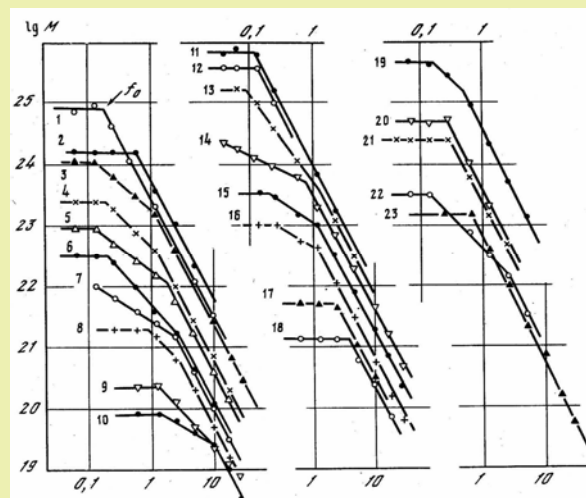
Maeda&Walter 1996

a set of coda magnitudes can be converted to an accurate source spectrum $\dot{M}_0(f)$



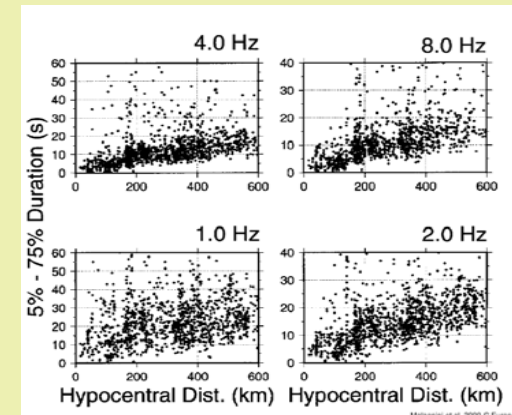
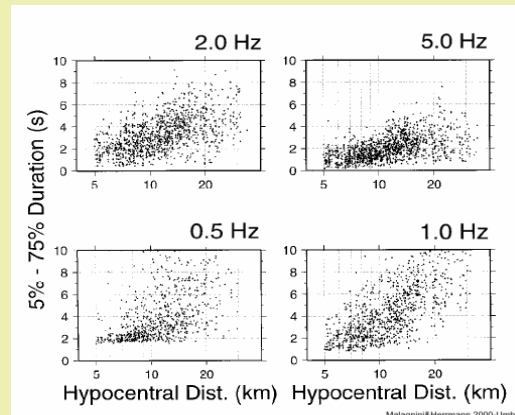
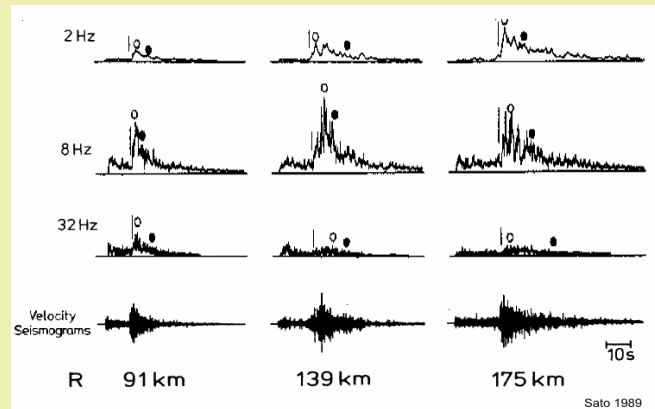
Maeda&Walter 1996

amplitude-based coda magnitude provides unsurpassed intrinsic accuracy:
 $\sigma(\text{single } \log_{10} A \text{ measurement}) = 0.05-0.1$,
 against 0.2-0.4 for usual magnitudes



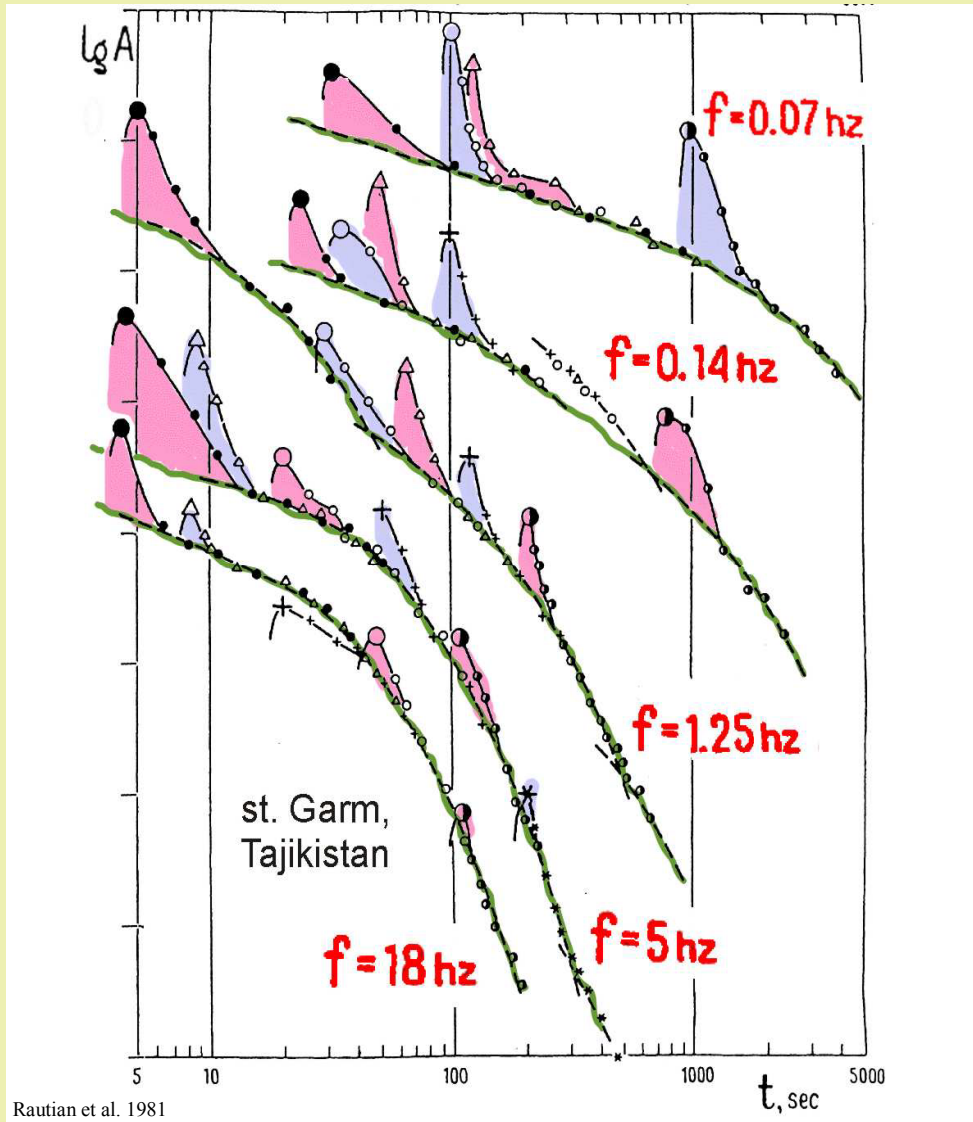
Rautian et al. 1981

Regional envelopes – body wave pulses



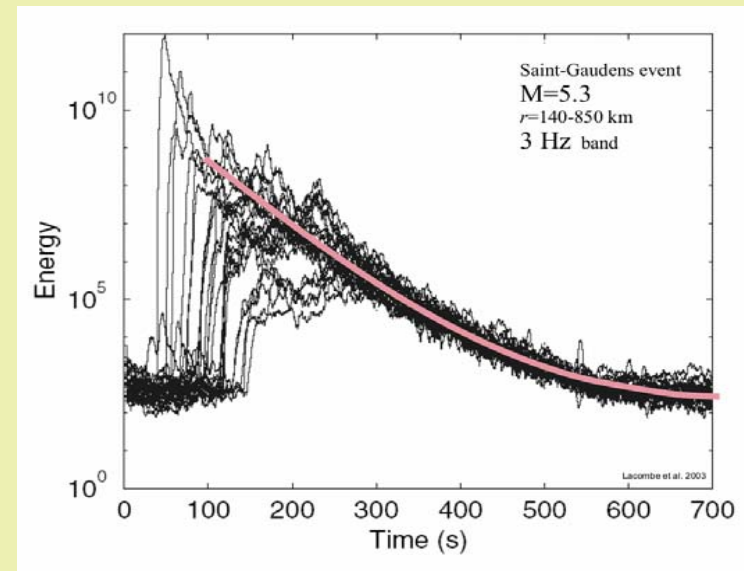
1. The duration of a body-wave group is difficult to parametrize and measure because of a very heavy coda tail. Different definitions can be based on:
 - ideally – mean delay of energy; in practice: onset-to-centroid or onset-to-peak time,
 - ideally - rms width of energy distribution, in practice: rms duration (“standard deviation”) of truncated data, or “interquantile range” of energy distribution in time, like 5%-75% range.
2. The duration of a body-wave group grows with hypocentral distance in the local (0-100km) and regional (0-600km) distance ranges. Pulse broadening is seen for oblique, near-horizontal and near-vertical rays. Lg over continental paths behaves differently, with saturation of duration.

Regional envelopes as a whole



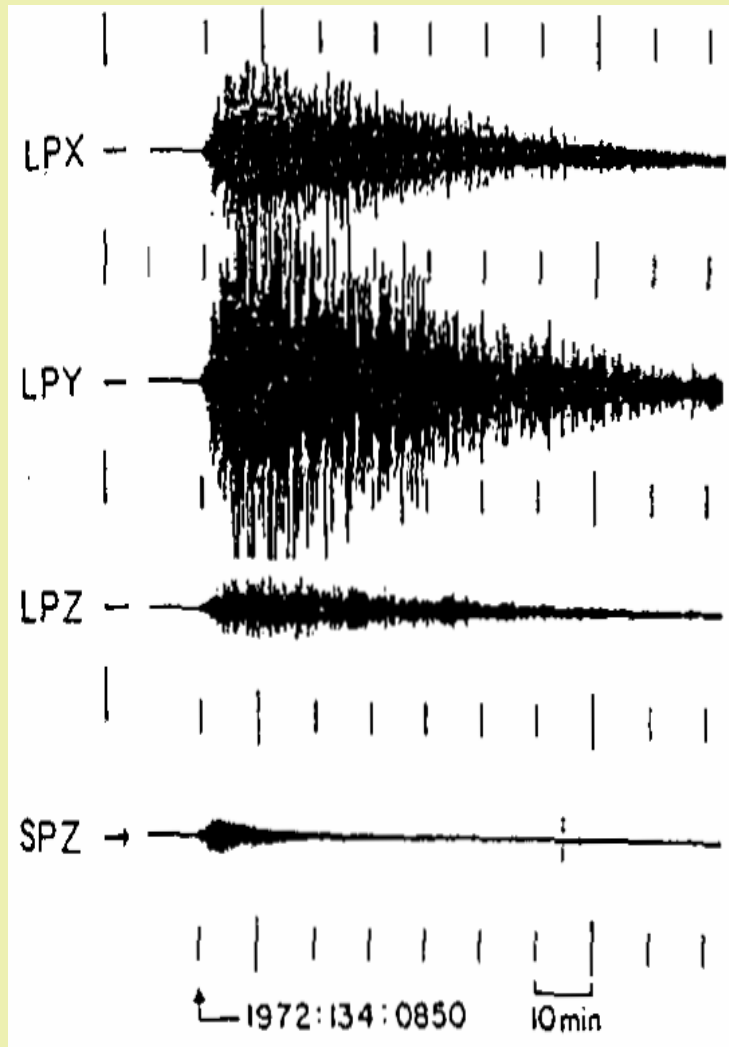
Over 20-30 to 500-1000 km distance range, S-wave group of increasing, medium-related duration is seen.

Typically S wave amplitudes are *above* coda asymptote.

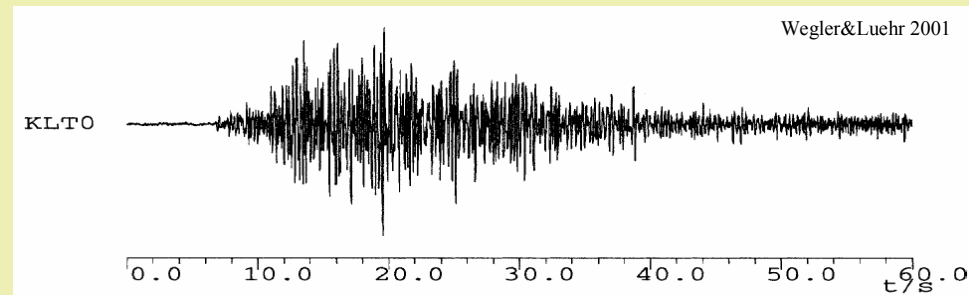


Diffusive envelopes – lunar, volcanic

lunar seismograms



B-type event on Merapi volcano



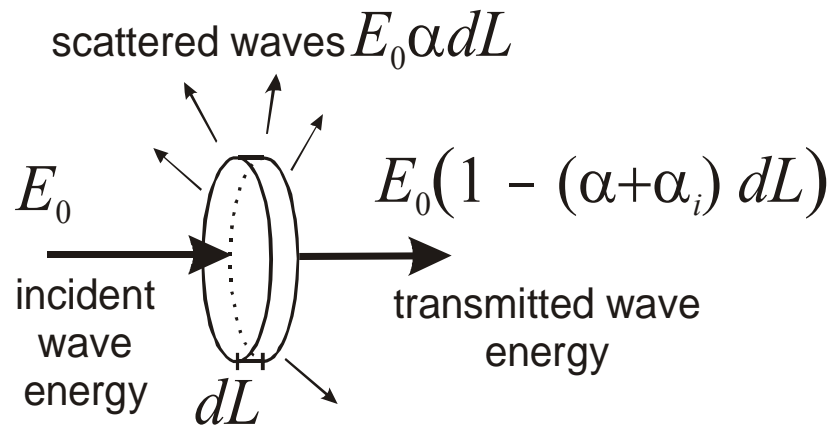
Spindle-like envelopes are characteristic for lunar seismograms and also for shallow events near volcanos (“Minakami B-type events”).

One sees very emergent onset, no direct body wave, no indications of S wave group. Coda is clear and stable.

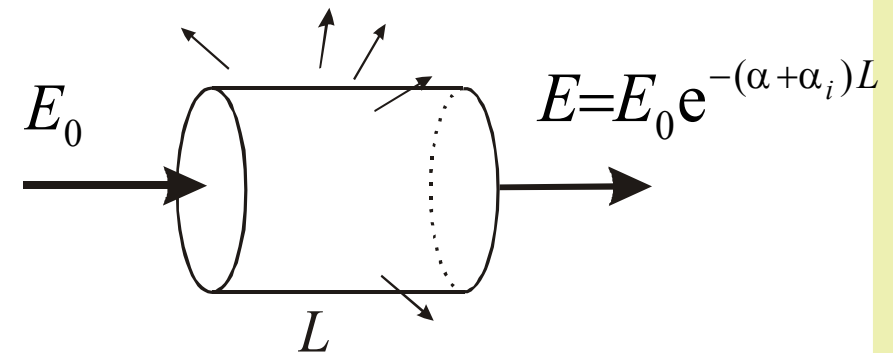
Such a picture is associated with wave energy diffusion in the medium (relatively very strong scattering).

(Contribution of source duration negligible)

Theory. Scattering coefficient or turbidity



integrating loss along incident ray:



α - scattering coefficient or turbidity (also α_s , also g)
 fractional loss of energy to scattering, per 1 km
 probability of scattering per 1 km
 units: km^{-1}

α_i - absorption coefficient

fractional intrinsic/inelastic loss, per 1 km

$\alpha_t = \alpha + \alpha_i$ - attenuation/extinction coefficient
 fractional *total* loss, per 1 km

Dimensionless quality factors Q are defined:

Q^{-1} = fractional loss per (wavelength/ 2π)
 so that

$$\alpha_s = 2\pi f / c Q_s \quad \alpha_i = 2\pi f / c Q_i \quad \alpha_t = 2\pi f / c Q_t$$

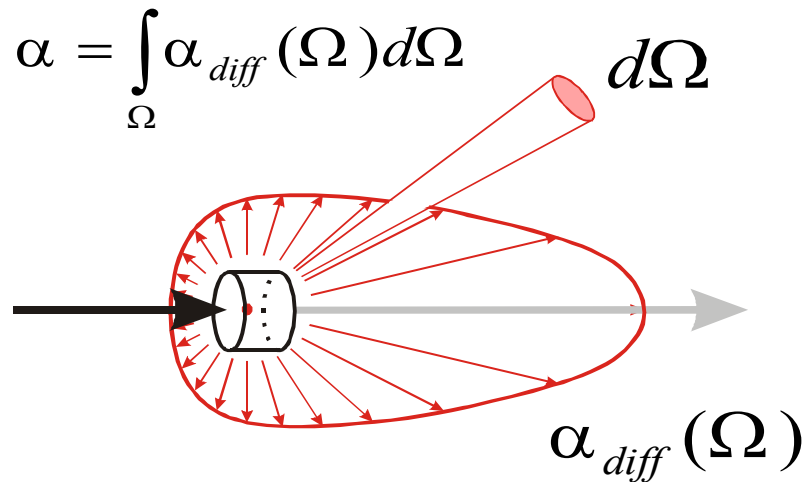
and: $1/Q_t = 1/Q_s + 1/Q_i$

for a beam of particles:

α is the probability of scattering per 1 km;
 hence:

Mean Free Path: $l = 1/\alpha$ [km]

Angular distribution of scattered energy. Phase function or indicatrix (1)

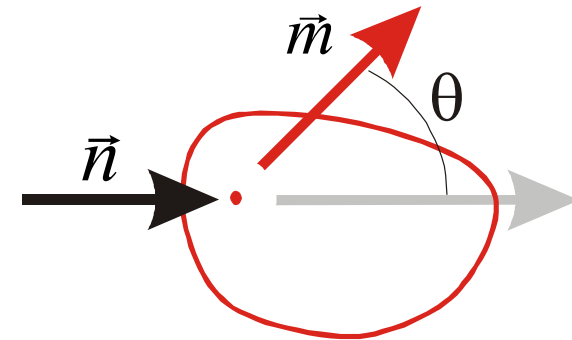


$\alpha_{diff}(\Omega)$ - differential scattering coefficient,
fractional scattering loss
per km per unit solid angle
(per steradian)

$\phi(\Omega) = \alpha_{diff}(\Omega) / \alpha$ - indicatrix or phase function

$$1 = \int_{\Omega} \phi(\Omega) d\Omega$$

$\phi(\Omega)$ can be treated as probability density
for a scattered particle to select a particular position
on a distant sphere around the scattering subvolume



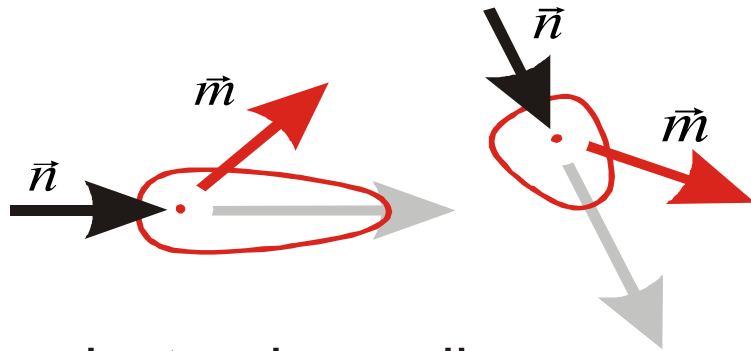
general case,

$$\phi() = \phi(\Omega_n, \Omega_m) \Rightarrow \phi(\vec{n}, \vec{m})$$

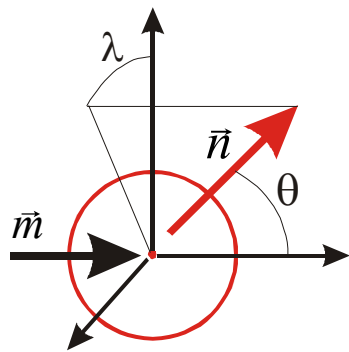
scattering angle :

$$\theta = \arccos(\vec{m} \cdot \vec{n})$$

Phase function (continued)

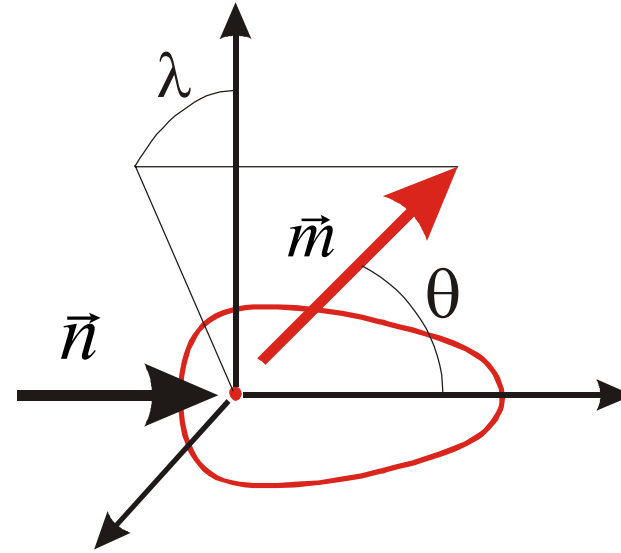


anisotropic-medium case
 (anisotropic w.r.t. N-E-Z reference,
 seems adequate e.g. for layered crust)



$$\phi(\vec{m}, \vec{n}) = \text{const} = \frac{1}{4\pi}$$

isotropic-medium and ray-isotropic case
 the simplest case



$$\phi(\vec{n}, \vec{m}) \Rightarrow \phi(\cos(\vec{n}, \vec{m})) \Rightarrow \phi(\theta)$$

non-isotropic,
 or anisotropic
 ("ray-anisotropic")
 case, axisymmetrical
 ("isotropic-medium" case, with statistically
 isotropic medium;
 no isotropy w.r.t. incident ray direction)

Equations of radiative transfer (stationary case)

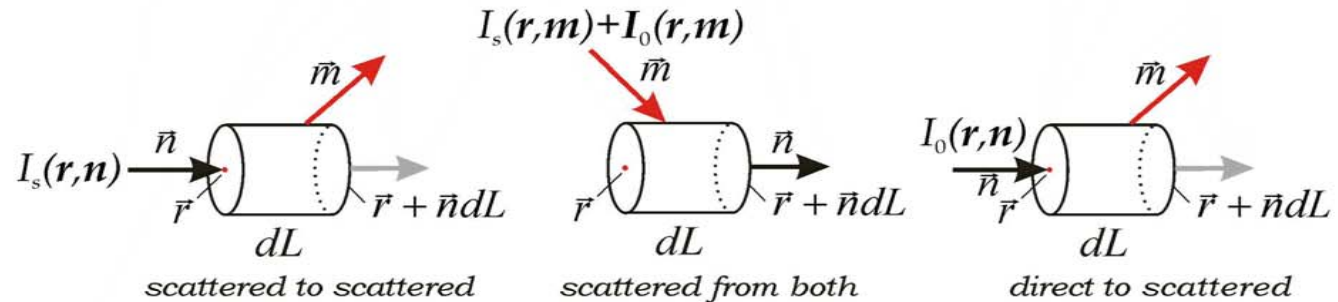
Define $I_s(\mathbf{r}, \mathbf{n})$

- scattered radiation intensity at \mathbf{r} along \mathbf{n}
 as: $dP_s = I_s(\mathbf{r}, \mathbf{n}) d\Omega_n$
 where dP_s is scattered wave power propagating from \mathbf{r} , along \mathbf{n} , into a cone with a solid angle $d\Omega_n$

Similarly, define $I_0(\mathbf{r}, \mathbf{n})$

- direct ("ballistic") radiation intensity at \mathbf{r} along \mathbf{n} (from a certain source).
 For the case of a point source, assume that a ray from it is along \mathbf{n} at \mathbf{r} .

(all this with respect to radiation in a certain frequency band Δf)



$$I_s(\mathbf{r} + \mathbf{n}dL, \mathbf{n}) - I_s(\mathbf{r}, \mathbf{n}) = -dI_1 + dI_2 = (\text{loss}) + (\text{gain})$$

$$\text{loss: } dI_1 = \alpha I_s(\mathbf{r}, \mathbf{n}) dL + \alpha_i I_s(\mathbf{r}, \mathbf{n}) dL \quad [\text{scatt. + intr.}],$$

(here α is the sum over all m !)

$$\text{gain: } dI_2 = \alpha \int_{4\pi} (I_s(\mathbf{r}, \mathbf{m}) + I_0(\mathbf{r}, \mathbf{m})) \phi(\mathbf{m}, \mathbf{n}) d\Omega_m$$

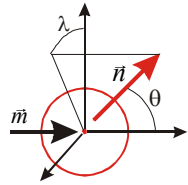
and similarly for I_0 , giving:

$$\frac{dI_s(\mathbf{r}, \mathbf{n})}{dL} = -\alpha I_s(\mathbf{r}, \mathbf{n}) - \alpha_i I_s(\mathbf{r}, \mathbf{n}) + \alpha \int_{4\pi} (I_s(\mathbf{r}, \mathbf{m}) + I_0(\mathbf{r}, \mathbf{m})) \phi(\mathbf{m}, \mathbf{n}) d\Omega_m$$

$$\frac{dI_0(\mathbf{r}, \mathbf{n})}{dL} = -\alpha I_0(\mathbf{r}, \mathbf{n}) - \alpha_i I_0(\mathbf{r}, \mathbf{n})$$

(in the non-stationary case, use $I_s = I_s(\mathbf{r}, t, \mathbf{n})$, and $\frac{d}{dL} = \mathbf{n} \cdot \nabla + \frac{1}{c} \frac{\partial}{\partial t}$ instead of $\frac{d}{dL}$)

Isotropic scattering case: general



$$\phi(\vec{m}, \vec{n}) = \text{const} = \frac{1}{4\pi}$$

isotropic-medium and ray-isotropic case
the simplest case

consider the simplest case:

- instant point source flashing at $t=0$,
- unit source energy
in the frequency band $(f-\Delta f, f+\Delta f)$;
- acoustic/scalar waves:
no conversion, no polarization
- isotropic scattering

DEFINITIONS

basic parameters:

r	source to receiver distance;
c	body wave speed (in applications, mostly S-wave speed);
$f, \Delta f$	wave frequency and bandwidth; $\omega = 2\pi f$
$\lambda = c/f$	wavelength
$k = 2\pi/\lambda = \omega/c$	wavenumber
$P(r, t)$	wave intensity in the same band (omnidirectional);
$P_c(t)$	coda intensity: $P(r, t) \rightarrow P_c(t)$ when $t \gg r/c$
l	mean free path
$t^* = l/c$,	mean free time
Q	quality factor due to scattering ($Q = \omega t^*$)

dimensionless / scaled parameters:

$\rho \equiv r/l$	scaled distance
$\tau \equiv cr/l = t/t^*$	scaled lapse time
$i(\rho, \tau), i_c(\tau)$	scaled scattered intensity (3D, use l^2 for 2D):

$$i(\rho, \tau) = \left(\frac{l^3}{c} \right) P(r, t)$$

$$i_c(\tau) \equiv i(0, \tau) \quad \text{scaled coda intensity:}$$

OMNIDIRECTIONAL WAVE INTENSITY

$$P_s(\mathbf{r}, t) = \int_{4\pi} I_s(\mathbf{r}, t, \mathbf{n}) d\Omega_n \quad \text{scattered}$$

$$P(\mathbf{r}, t) = \int_{4\pi} (I_0(\mathbf{r}, t, \mathbf{n}) + I_s(\mathbf{r}, t, \mathbf{n})) d\Omega_n \quad \text{total}$$

Isotropic scattering case: SIS

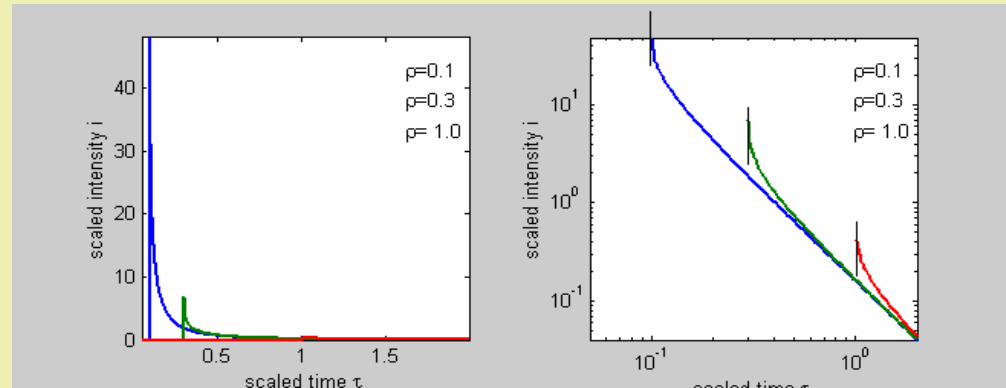
$$\rho \ll 1, \tau \ll 1$$

Single (isotropic) scattering model - SIS

(single= Born approximation):,

$$i^{\text{SIS}}(\rho, \tau) = \frac{1}{4\pi\rho\tau} \ln\left(\frac{\tau + \rho}{\tau - \rho}\right)$$

$$i_c^{\text{SIS}}(\tau) = \frac{1}{2\pi\tau^2}$$



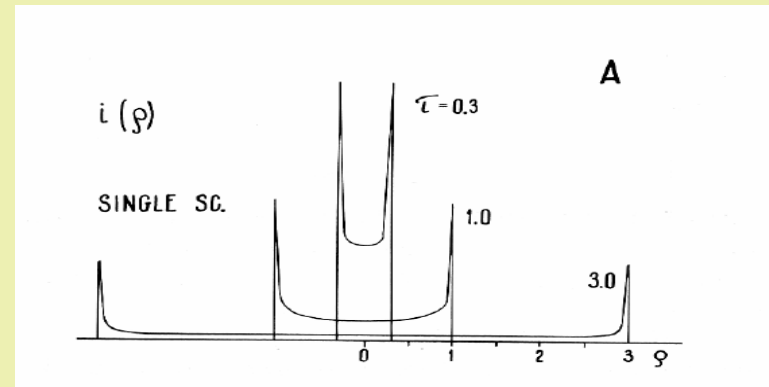
Main properties:

A. “positive” [fit regional waveforms]

1. Clear coda
2. Clear coda asymptote
3. Pulse envelope approaches coda asymptote from above

B. “negative” [contradict regional waveforms]

1. Spike-like arrival, no pulse broadening with distance
2. Inaccurate at $\rho \cong 1$ or more



“Coda-Q” determination:

fit the observed coda shape selecting Q_c in the equation

$$I_c^{\text{SIS}}(t) = \frac{\exp(-2\pi ft / Q_c)}{2\pi c l t^2}$$

Isotropic scattering case: diffusion approximation

$\tau \gg 1$, any ρ

Diffusion isotropic scattering model – DIS

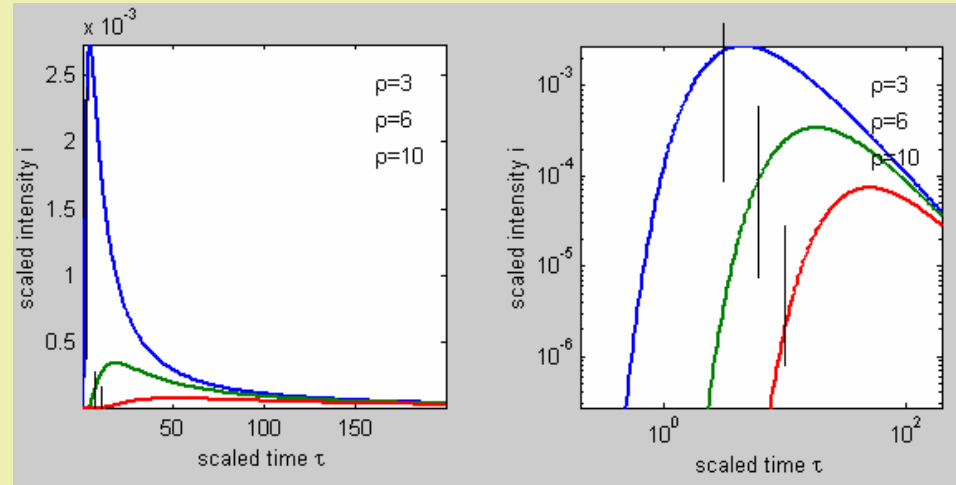
$$i^{\text{DIS}}(\rho, \tau) = \frac{1}{(4/3\pi\tau)^{3/2}} \exp\left(-\frac{\rho^2}{4/3\tau}\right)$$

$$i_c^{\text{DIS}}(\tau) = \frac{1}{(4/3\pi\tau)^{3/2}}$$

the solution of parabolic/diffusion equation
for wave energy density $E(\mathbf{r}, t) = P(\mathbf{r}, t)/c$:

$$\partial E / \partial t = D \nabla^2 E$$

where $D = lc/3$ in 3-dim.case (or $lc/2$ in 2D)



Main properties:

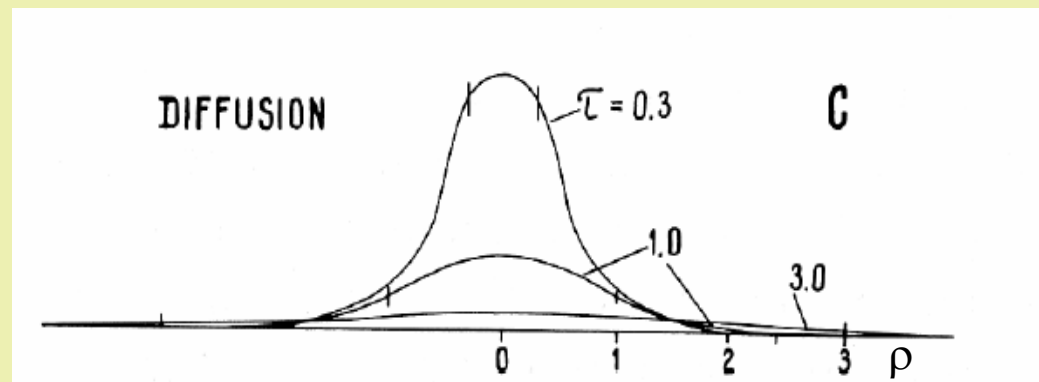
A. “positive” [fit regional waveforms]

1. Clear coda, clear coda asymptote
2. “Pulse” broadens with distance

B. “negative” [contradict regional waveforms]

1. “Pulse” envelope approaches coda asymptote from below
2. Weak arrival
3. “Pulse” is too long
4. In space, energy concentrates around the source
5. Bad model at $\rho \cong 2$ or less

C. conclusion: Can fit lunar and volcanic data but not regional waveforms



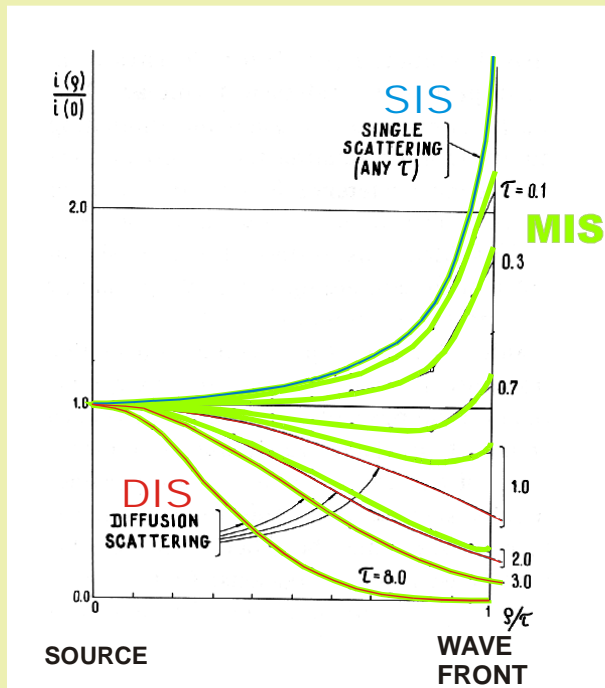
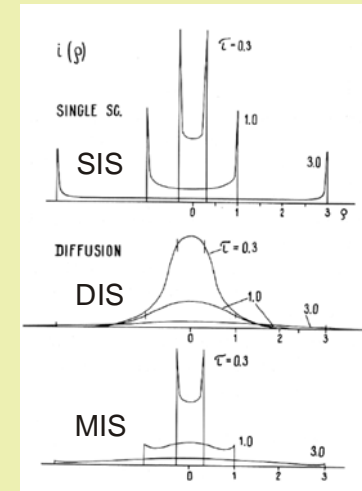
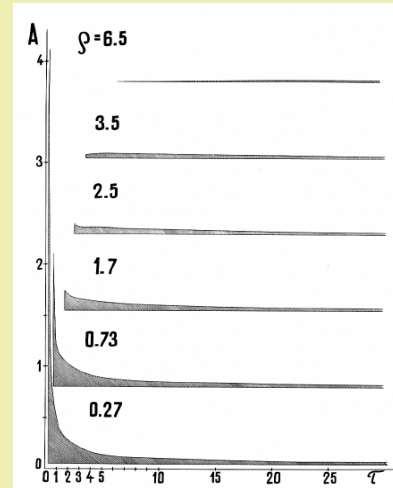
Isotropic scattering case: multiple

Multiple isotropic scattering model - MIS

any τ , any ρ

$i^{\text{MIS}}(\rho, \tau) \lll$ Numerical MC model (Gusev & Abubakirov 1987) \ggg
 Analytical series representation (Zeng et al. 1991)

$$i_c^{\text{MIS}}(\rho, \tau) \cong \frac{1}{2\pi\tau^2} \left[1 + \left(\frac{27}{16\pi} \tau \right)^x \right]^{-1/2x} ; x = 1.10 \quad (\text{Abubakirov \& Gusev 1990})$$



Main properties:

A. “*positive*” [fit regional waveforms]

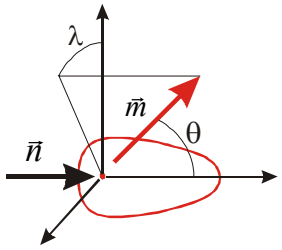
1. Clear coda & coda asymptote

B. “*negative*” [contradict regional waveforms]

1. Spike-like arrival (or very long train):

no realistic pulse broadening with distance

Multiple non – isotropic scattering



$\phi(\vec{n}, \vec{m}) \Rightarrow \phi(\cos(\vec{n}, \vec{m})) \Rightarrow \phi(\theta)$
 non-isotropic,
 or anisotropic
 ("ray-anisotropic")
 case, axisymmetrical
 ("isotropic-medium" case, with statistically
 isotropic medium;
 no isotropy w.r.t. incident ray direction)

Instead of a single $l \equiv$ MFP in the isotropic case,
 two characteristic lengths:
 (1) l_n - "non-isotropic",
 "true" MFP,
 (2) l - transport MFP,
 defined through diffusion asymptotics ($t \rightarrow \infty$) as
 $l = 3D/c$ (in 3-dim.case)

dimensionless / scaled parameters:

$\rho \equiv r/l$ scaled distance ("transport")
 $\tau \equiv cr/l = t/t^*$ scaled lapse time ("transport")
 $\rho_n \equiv r/l_n$ scaled distance ("common, true")
 $\tau_n \equiv crt/l_n = t/t_n^*$ scaled lapse time ("common, true")
 $i(\rho, \tau), i_c(\tau)$ scaled scattered intensity
 (3D, use l^2 for 2D):

$$i(\rho, \tau) = i\left(\frac{r}{l}, \frac{t}{t^*}\right) = \left(\frac{l^3}{c}\right) P(r, t)$$

scaled coda intensity:

$$i_c(\tau) \equiv i(0, \tau)$$

MORE DEFINITIONS

basic parameters:

$l, t^* = l/c$, redefined as transport mean free path,
 and transport mean free time
 (compatible to previous definition)

$l_n, t_n^* = l_n/c$, (common) mean free path,
 and mean free time

Q transport quality factor due to scattering
 ($Q = \omega t^* = 2\pi l / \lambda$);

Q_n (common) quality factor due to scattering
 ($Q_n = \omega t_n^* = 2\pi l_n / \lambda$);

KEY FORMULA FOR TRANSPORT MFP

$$l = \frac{l_n}{1 - \langle \cos \theta \rangle}$$

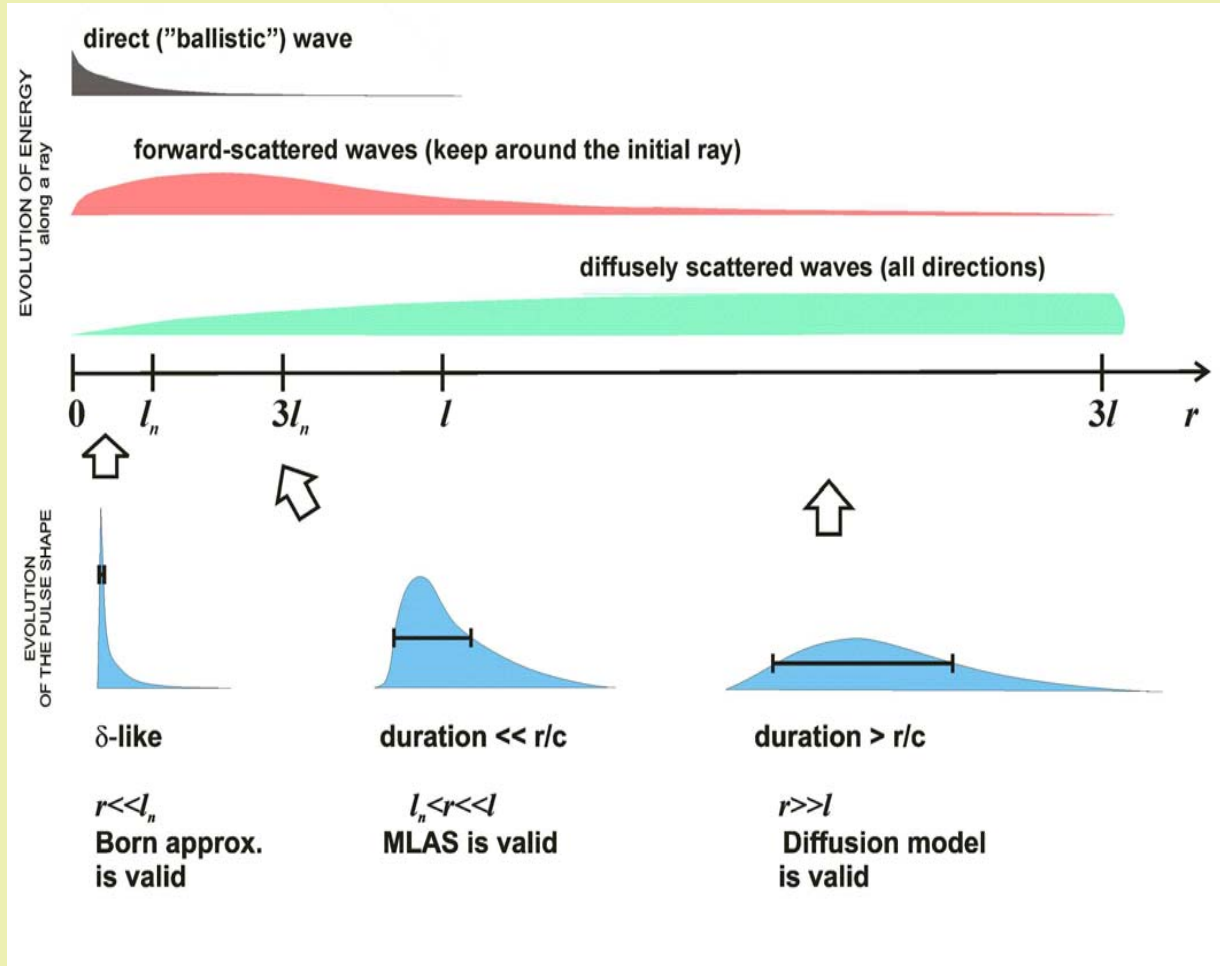
where

$$\langle \cos \theta \rangle = \frac{1}{4\pi} \int \phi(\Omega) \cos \theta d\Omega =$$

$$= \int_0^\pi \int_0^{2\pi} \phi'(\theta) \cos \theta \sin \theta d\lambda d\theta$$

Typical value for the Earth's lithosphere: $l \equiv$ MFP = 100 km, so for typical local/regional observations: $\rho = 0.3 - 2$

Multiple low-angle scattering



FORWARD-ENHANCED
(NARROW) PHASE FUNCTION

$$\langle \theta^2 \rangle \ll 1$$

$$l_n/l = 1 - \langle \cos \theta \rangle \approx \langle \theta^2 \rangle / 2 \ll 1$$



DEFINITIONS

OF scattering- Q :
standard:

$$Q = 2\pi l_n / \lambda$$

(direct \rightarrow

\rightarrow forward-scattered)

in seismology, in practice

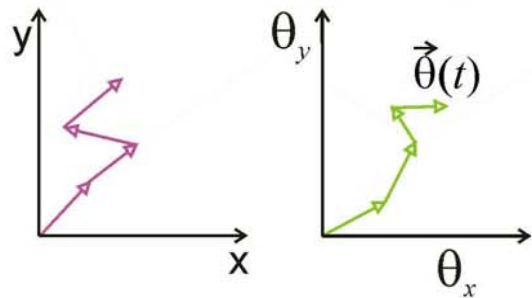
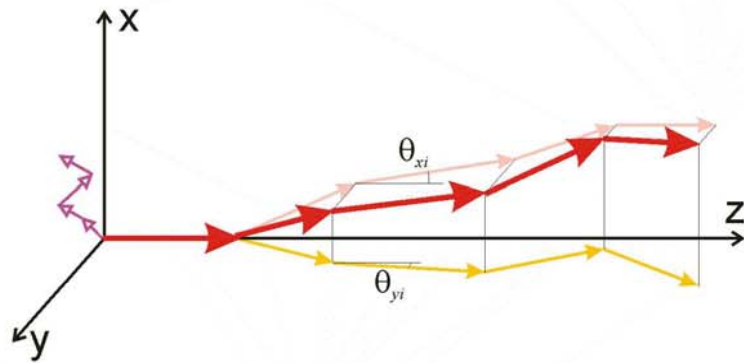
$$Q = 2\pi l / \lambda$$

(direct + forward-scattered \rightarrow

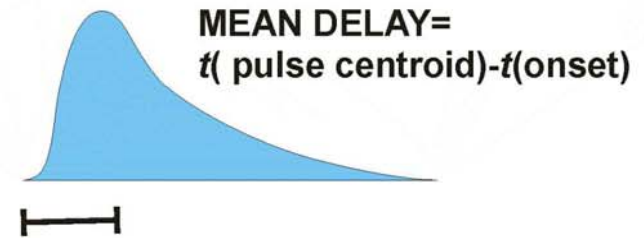
\rightarrow diffusely-scattered)

[related to the habit to integrate
entire "body-wave group"
as *direct wave*]

Multiple low-angle scattering(2)



Assume
 $\langle \theta_{xi}^2 \rangle < \infty$
 $\langle \theta_{yi}^2 \rangle < \infty$
 then $\vec{\theta}(t)$ is a
 Brownian motion



$$\langle T \rangle = \frac{\int_{t_d}^{\infty} (t - t_d) E(t) dt}{\int_{t_d}^{\infty} E(t) dt}$$

$$\langle T \rangle = \frac{r^2}{6cl}$$

$$\tau_a = \frac{\langle T \rangle}{t^*} = \frac{1}{6} \rho^2$$

Multiple non-isotropic scattering – simulation

Monte-Carlo simulation:
the standard technique
to solve real
radiative transport problems.
No ready analytic solution
exists
for multiple non-isotropic
scattering
even in the case of uniform-
space geometry
and isotropic-medium phase
function

EXAMPLE

2D, $\tau=0.7$, $N=500$

source:

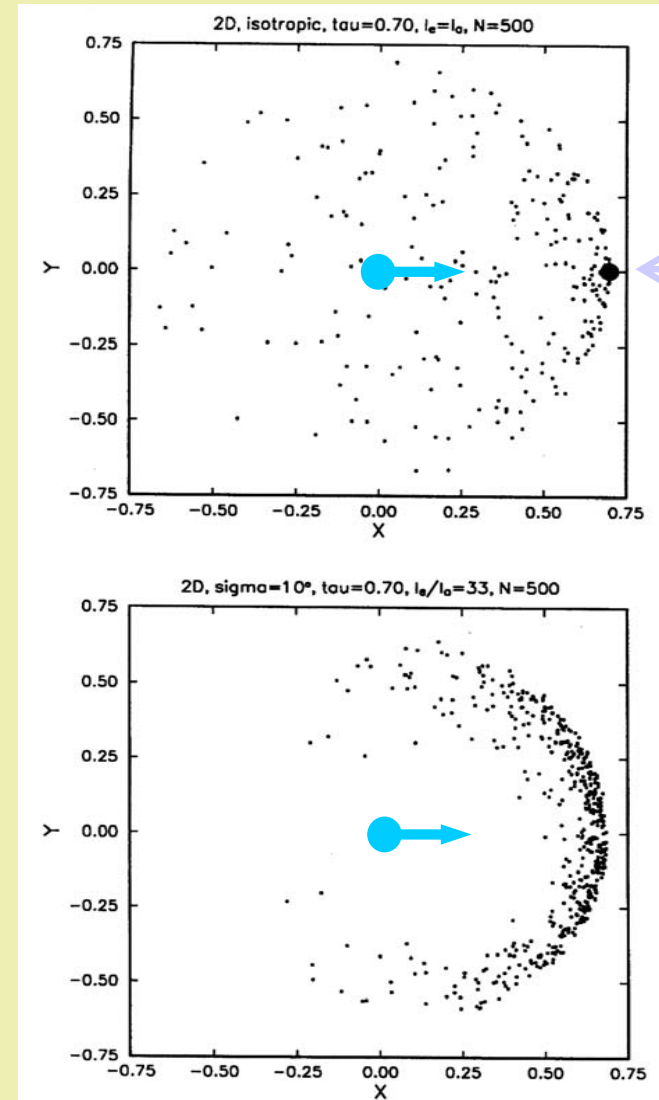
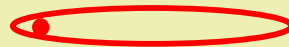
needle-like radiation pattern
along +x



phase function:
isotropic

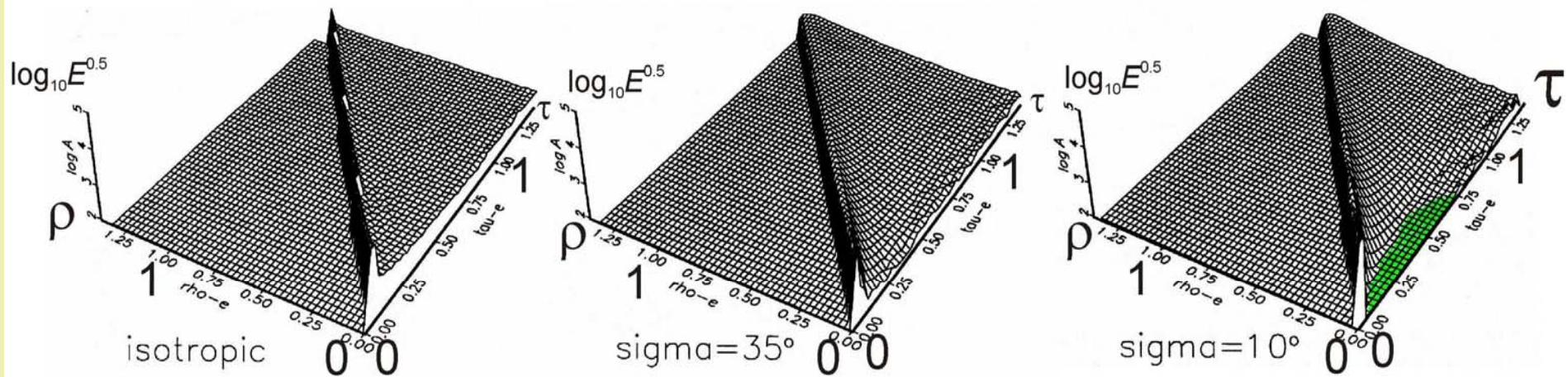


phase function:
 $(\langle\theta^2\rangle)^{0.5}=\sigma=10^\circ$



ballistic/direct component

Multiple non-isotropic scattering – simulated envelopes



Isotropic scattering case:

spike-like energy pulse – no broadening,
completely unrealistic

well-formed, monotonous,
believable coda

Moderately elongated phase function ($\sigma=35^\circ$):

acceptably broadening energy pulse

no minimum in coda,
marginally acceptable

Narrow phase function ($\sigma=10^\circ$):

well-formed, broadening energy pulse

coda with minimum,
completely unrealistic

CONCLUSION: Both isotropic-scattering and MLAS models do not work.
Real phase function must be moderately elongated

Which parameter specifies the scattering properties of the Earth's medium?

Three modes of analysis of observed signals are used to extract scattering properties of the Earth's medium:

(1) The ratio of coda amplitude to S-wave pulse amplitude

gives l - transport MFP

{traditionally, viewed as
“back-scattering MFP”
or “isotropic-scattering MFP”}

[in an improved form, works as a part of MLTWA]

(2) The rate of S-wave pulse energy attenuation with distance

gives $Q_{total} \Rightarrow l$ - transport MFP

{traditionally, the “scattering part” of Q_{total}^{-1} is
treated as “the” scattering Q^{-1} and
associated with “isotropic-scattering MFP”}

[in a modified form, works as a part of MLTWA]

(3) Pulse broadening rate with distance

gives l - transport MFP

**No technique has been proposed in
seismology to determine l_n - true MFP**

**and there are theoretical obstacles that
complicate such a determination**

**A certain confusion is produced by using
isotropic scattering model in the
interpretation of observations**

**whereas in the Earth, the phase function is
definitively forward-enhanced**

**In reality, most techniques that aimed at
determination of MFP (or scattering Q), yield
transport MFP**

CONCLUSION:

**one can continue to use the usual
“seismological” scattering- Q parameter**

**but should keep in mind that it essentially
related to transport MFP,
and *not* to true MFP**

Random inhomogeneity field and phase function

Random medium –
the simplest case
(for the Earth, essentially, each
assumption is an oversimplification)

Acoustic/scalar waves:

$$c(\mathbf{x}) = c_0 (1 + \varepsilon'(\mathbf{x}))$$

Weak inhomogeneity:

$$\varepsilon'(\mathbf{x}) \ll 1$$

Gaussian inhomogeneity -
can be described by ACF:

$$\text{Cov}(\varepsilon'(\mathbf{x}), \varepsilon'(\mathbf{y}))$$

Stationary inhomogeneity:

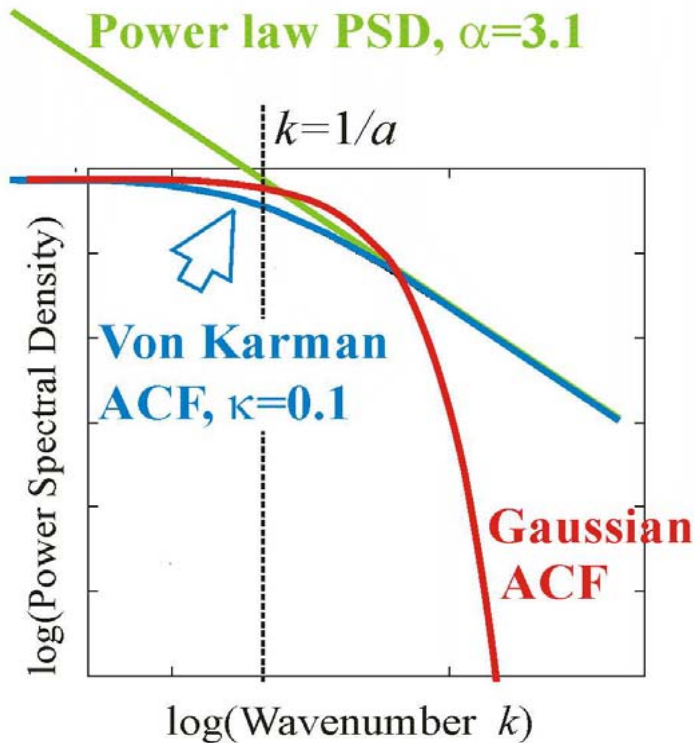
$$\begin{aligned} \text{Cov}(\varepsilon'(\mathbf{y}), \varepsilon'(\mathbf{y} + \mathbf{x})) &= \\ &= \sigma_\varepsilon^2 R'(\mathbf{x}) \end{aligned}$$

Isotropic inhomogeneity:

$$\begin{aligned} \text{Cov}(\varepsilon'(\mathbf{y}), \varepsilon'(\mathbf{y} + \mathbf{x})) &= \\ &= \sigma_\varepsilon^2 R'(\mathbf{x}) = \sigma_\varepsilon^2 R(|\mathbf{x}|) = \\ &= \sigma_\varepsilon^2 R'(r) \end{aligned}$$

Case	ACF	POWER SPECTRUM $k' = \mathbf{k}' $ is related to FT[$\varepsilon(\mathbf{x})$ of <i>medium</i>]	PHASE FUNCTION $k = \mathbf{k} = \omega/c$ is related to <i>propagating waves</i>
General	$R(r)$	$\tilde{R}(k')$	$\phi(\theta) \propto k^4 \tilde{R}(2k \sin(\theta/2))$
Gaussian ACF	$\exp(-r^2 / a^2)$	$\propto \exp(-(ka)^2 / 4)$	$\phi(\theta) = \frac{\exp((\cos \theta - 1)/\sigma^2)}{2\pi\sigma^2(1 - \exp(-2/\sigma^2))}$ where $\sigma^2 = 2/(ka)^2$
self-affine	diverges at $r = \infty$	$k^{-\alpha}$	$\propto (\sin(\theta/2))^{-\alpha}$ diverges at $\theta = 0$
Von Karman	$\propto \left(\frac{r}{a}\right)^\kappa K_\kappa\left(\frac{r}{a}\right)$	$\propto \frac{1}{(1 + a^2 k^2)^{\kappa+3/2}}$ $\approx k^{-(2\kappa+3)}$ when $k \gg 1/a$	$\propto k^2 (1 + 4a^2 k^2 \sin^2(\theta/2))^{-(\kappa+3/2)}$ $\approx \sin(\theta/2)^{-(2\kappa+3)}$ as $k \gg 1/a$ (i.e. at not very small θ)

Random inhomogeneity field: models



The case of self-similar inhomogeneity:

$$\alpha=3$$

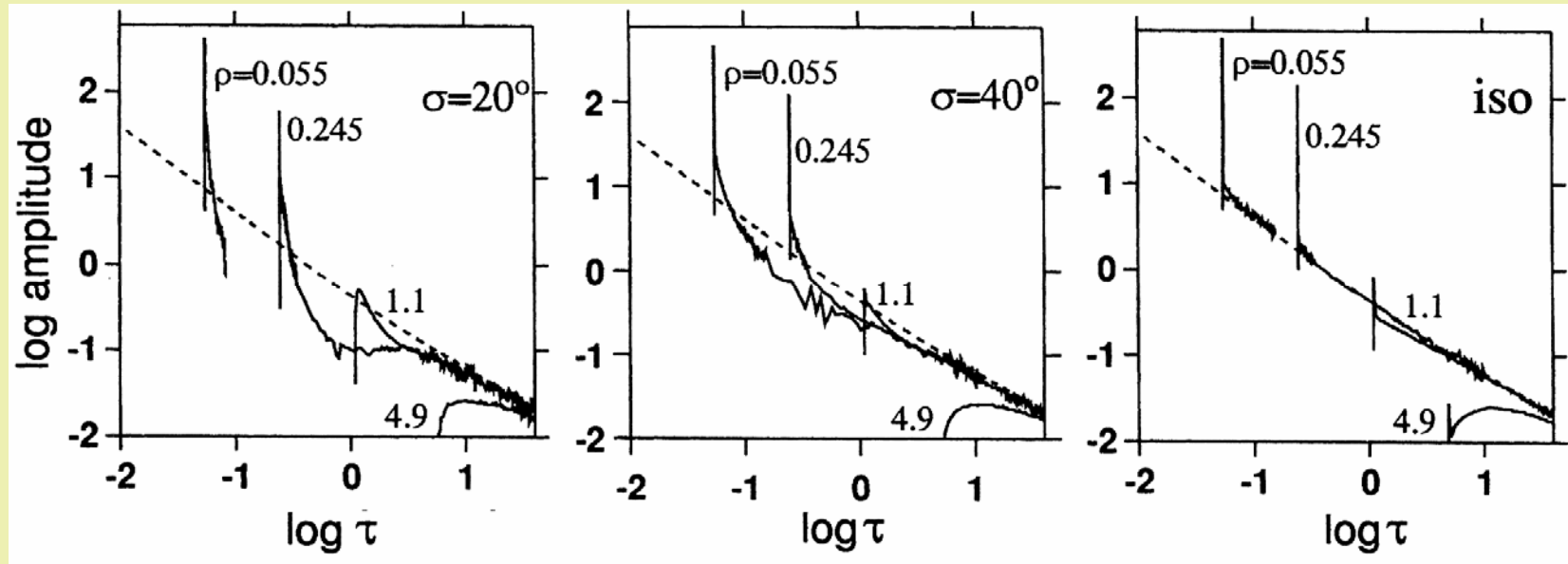
$$\kappa=0$$

Case	Properties of phase function $\phi(\theta)$ and power spectral density (PSD)
Gaussian ACF:	<p>$\phi(\theta)$: The angular width is strongly frequency-dependent:</p> $\sigma = 2^{0.5}/ka.$ <p>PSD: Abrupt high-wavenumber cutoff</p>
Self-affine case, power-law PSD:	<p>$\phi(\theta)$: Frequency-independent shape for all θ</p> <p>PSD: <i>Non-integrable</i> (in practical calculation, PSD can be truncated at small k)</p>
Von Karman ACF	<p>$\phi(\theta)$: Through selecting a sufficiently large value of a, one can provide the frequency-independent behavior of $\phi(\theta)$ for almost all θ, except for very small θ ($<1/ka$).</p> <p>PSD: Integrable.</p>

Models of random inhomogeneity field vs. reality

Case	comments
Gaussian ACF:	<p>Qualitatively unacceptable model.</p> <p>The strong frequency dependence ($1/k \rightarrow 1/f$) of the width σ of phase function makes impossible to match the requirement: $\sigma \approx 25-40^\circ$ - simultaneously for many frequency bands.</p>
Self-affine case, power-law PSD or Von Karman-ACF case with large α	<p>Qualitatively acceptable model.</p> <p>The frequency-independent shape of phase function for all or almost all angles enables one to fit the qualitative behavior of envelopes simultaneously for many frequency bands.</p> <p>[rough ranges for parameters: $\alpha=3.2-4$; $\kappa=0.1-0.5$]</p>

Simulated envelopes: Gaussian-ACF case



$\sigma=20^\circ$

- (1) gap instead of coda
- (2) pulse broadens with distance

$\sigma=40^\circ$

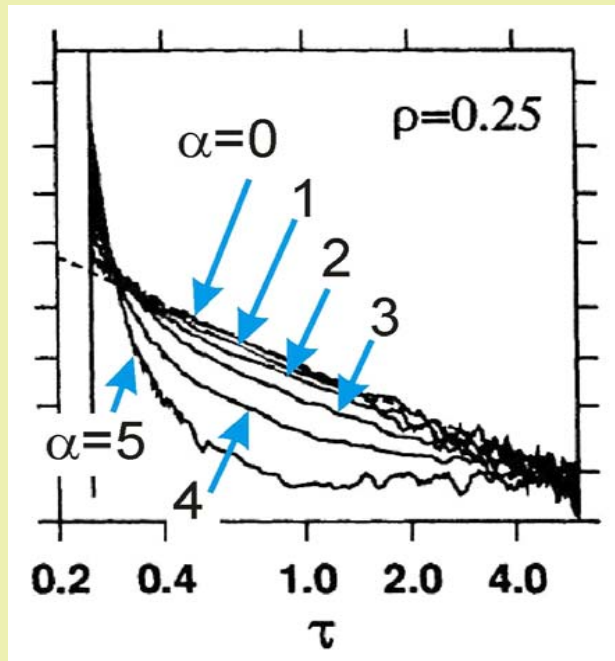
- (1) acceptable coda, note that its level is below that for isotropic-scattering case
- (2) spike instead of pulse up to $\rho \approx 1.5$

isotropic scattering ($\sigma=\infty$)

- (1) “perfect” coda
- (2) no pulse broadening at all

The interval estimate for σ , namely $\sigma = 20-40^\circ$, is attained, but it works for a single frequency band only! Gaussian-ACF model is mostly of instructional interest!

Simulated envelopes: self-affine case

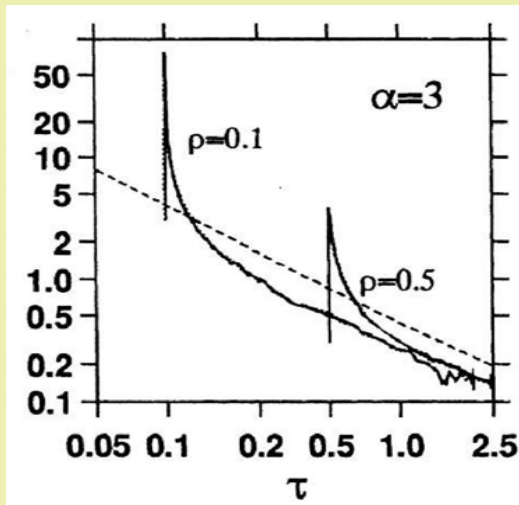


$\alpha=3$

- (1) quite acceptable coda shape
- (2) slightly too abrupt pulse onset

$\alpha=4$

- (1) early coda somewhat too low
- (2) acceptable pulse shape



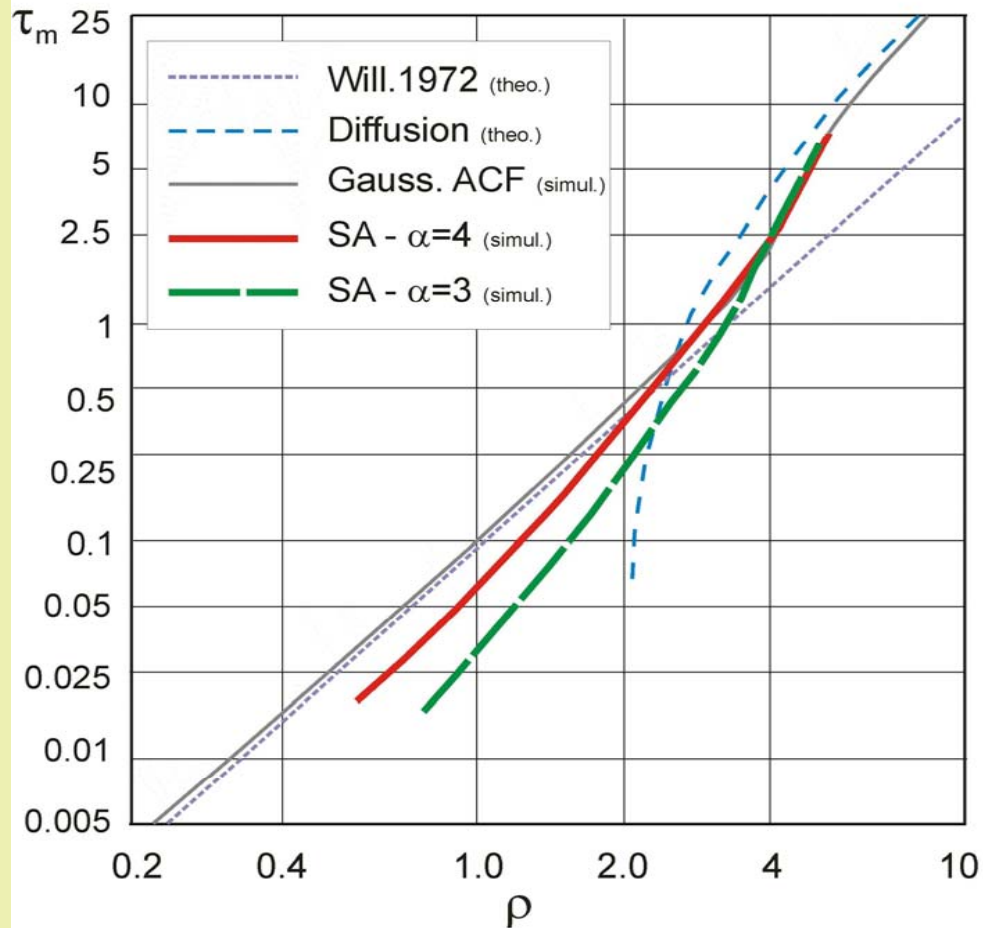
CONCLUSION

- (1) **Self-similar random inhomogeneity with $\alpha=3.2-4$ is a reasonable starting model for the lithosphere**
- (2) **Coda levels are systematically somewhat lower w.r.t. those of the isotropic scattering model ($\alpha=0$)**



Duration of simulated envelopes

Scaled onset-to-peak delay time τ_m
vs. scaled distance ρ



Gaussian-ACF case,
narrow phase function:

$$\tau_m = 0.091 \rho^2$$

on condition $\rho \ll 1$
(Williamson 1972)

Onset-to-peak delay for a realistic self-similar medium is significantly *smaller* than for the Gaussian-ACF medium.

When the α parameter can be specified or assumed, one can use the results of Monte-Carlo simulation to derive l from the observed duration trend.

Ways for inversion for scattering/attenuation parameters (body waves)

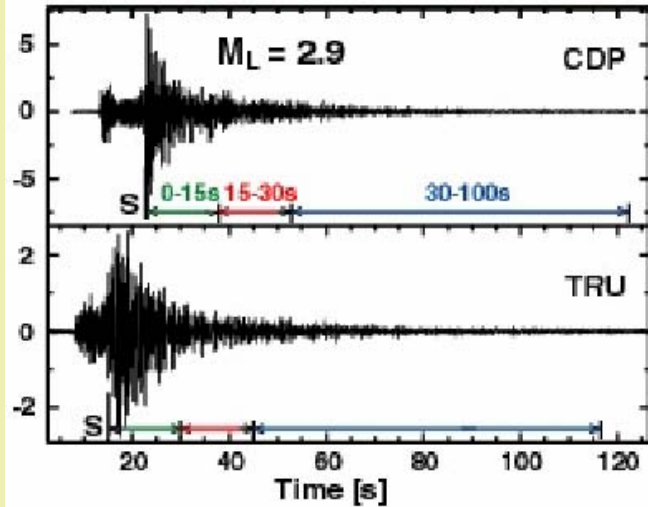
approach	comment
<p>A. Total attenuation</p> <p>$Q^{-1}_{total} [=Q^{-1}_{scattering} + Q^{-1}_{intrinsic}]$</p> <p>from body wave Fourier <i>spectra</i>.</p> <p>A1. From spectra as is – one (or more) events at many stations.</p> <p>A2. From spectra normalized to coda power at one or more stations</p>	<p>Efficient descriptive approach, valid for eventual synthetics.</p> <p>Results physically not transparent.</p> <p>Systematic, consistent selection of the data window difficult.</p> <p>Using coda normalization significantly reduces noise.</p>
<p>B. Total attenuation Q^{-1}_{total} from body wave <i>amplitudes</i>, raw or coda-normalized</p>	<p>Generally, outdated approach. Q^{-1}_{total} estimates often biased (because of variable, distance-dependent duration of the body wave group).</p>
<p>C. Separately $Q^{-1}_{scattering}$ <i>and</i> $Q^{-1}_{intrinsic}$ [add up to Q^{-1}_{total}] assuming isotropic scattering in uniform random medium.</p> <p>C1. By MLTWA (Multiple Lapse-Time Window Analysis) method</p> <p>C2. From Pulse-energy to coda-power ratio at the same propagation time.</p>	<p>Consistent separate estimates of $Q^{-1}_{scattering}$ and $Q^{-1}_{intrinsic}$.</p> <p>Results may be significantly model-dependent</p>

Ways for inversion for scattering/attenuation parameters (body waves) (2)

approach	comment
D. Only $Q^{-1}_{\text{scattering}}$ from body-wave pulse broadening.	Results may be model-dependent
E. Only $Q^{-1}_{\text{intrinsic}}$ from $\kappa(r)$ (κ in $A/A_0 = \exp(-\pi\kappa f)$)	Efficient but works only for frequency-independent component of attenuation. May be biased by effects of source spectra
F. Determination of “coda Q”	The approach assumes single isotropic scattering i.e. an unrealistic model, and cannot yield reliable results; but supported by a number of empirical parallels between Q_{total} and coda Q. Empirical coda Q is often lapse-time dependent, but other Q measures may behave similarly.

MLTWA (after Fehler 2003)

Integrate Energy in Windows whose Times are Referenced to S-wave Arrival Time

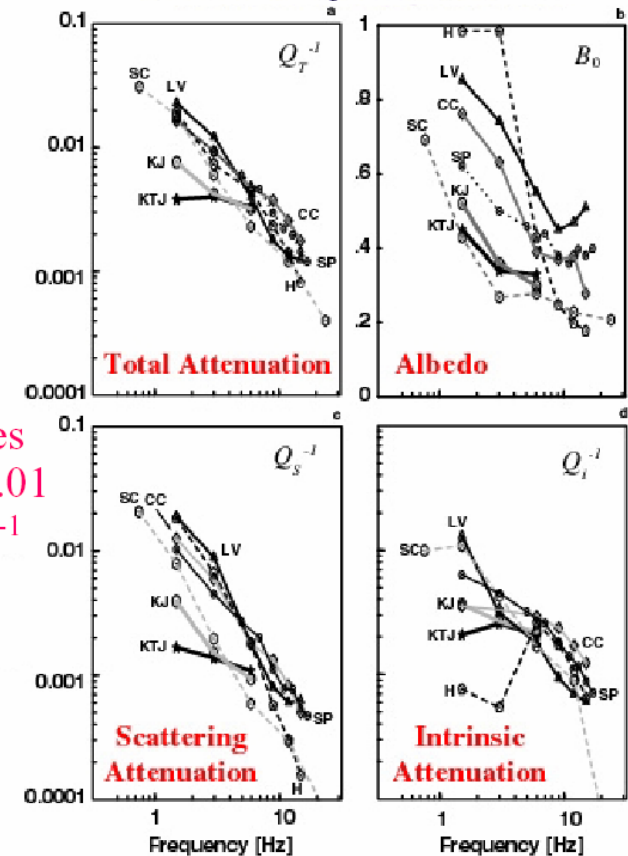


$$EI_1(f)_{kj} = \rho_0 \int_0^{15s} |\dot{u}_{kj}^S(t;f)|^2 dt,$$

$$EI_2(f)_{kj} = \rho_0 \int_{15s}^{30s} |\dot{u}_{kj}^S(t;f)|^2 dt,$$

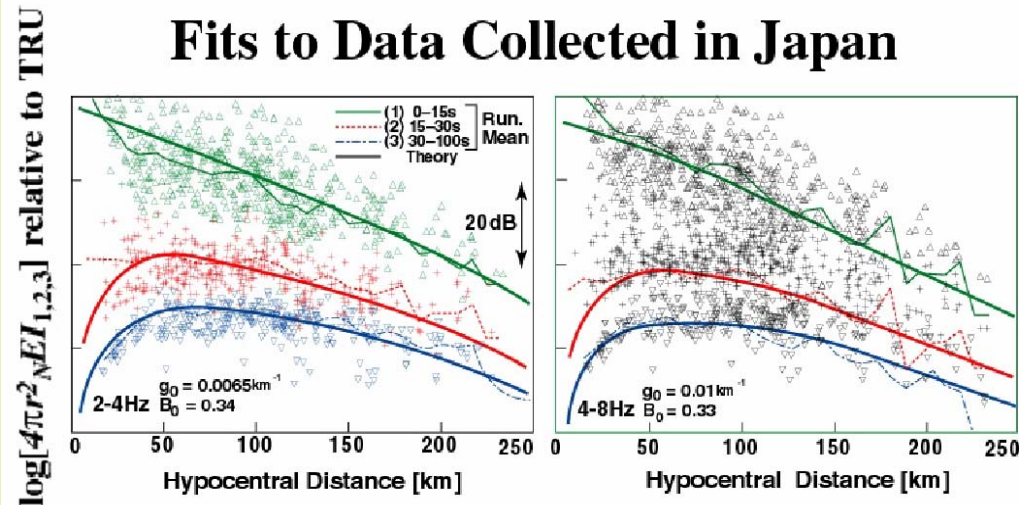
$$EI_3(f)_{kj} = \rho_0 \int_{30s}^{100s} |\dot{u}_{kj}^S(t;f)|^2 dt$$

Results Obtained with Multiple Lapse-Time Window Analysis Method

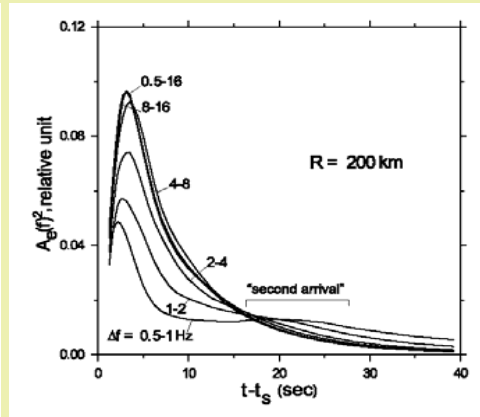
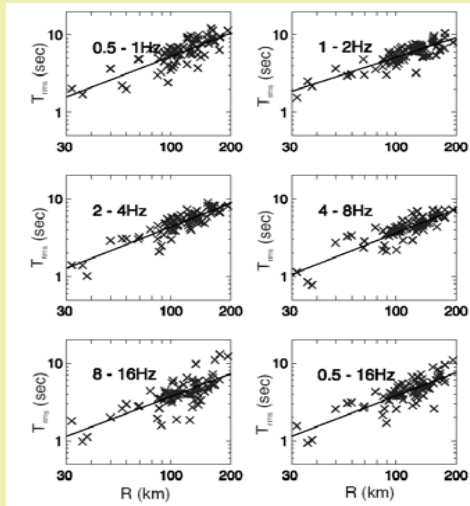


Gives $g_0 \sim 0.01 \text{ km}^{-1}$

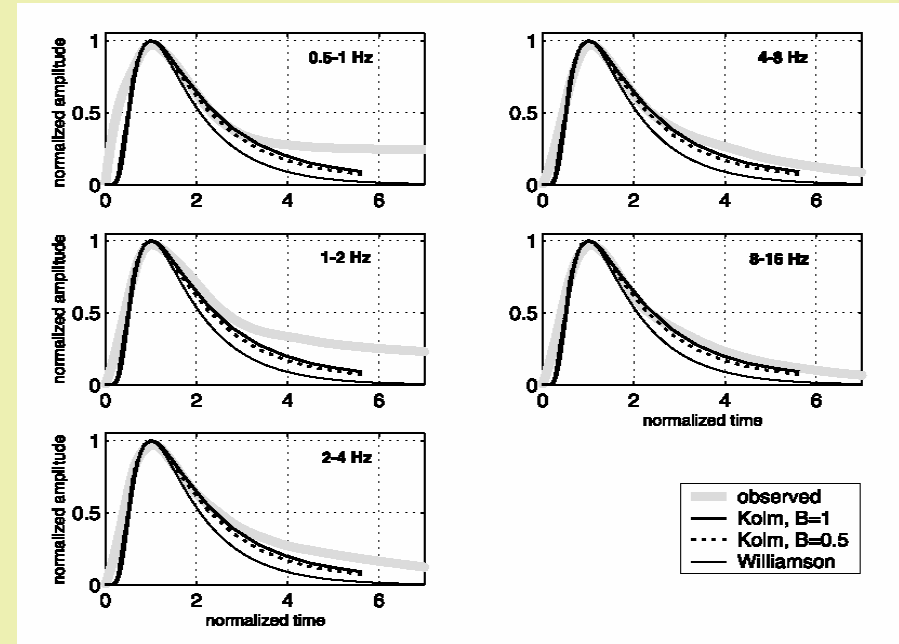
Fits to Data Collected in Japan



Scattering parameters from pulse duration vs distance trend



pulse shapes scaled along t axis, thus reduced to a fixed distance



RMS duration of S-wave group for sta. PET grows as $r^{-1.0}$ indicating strongly distance-dependent scattering Q .

To determine MFP, onset-to-peak delays are used.

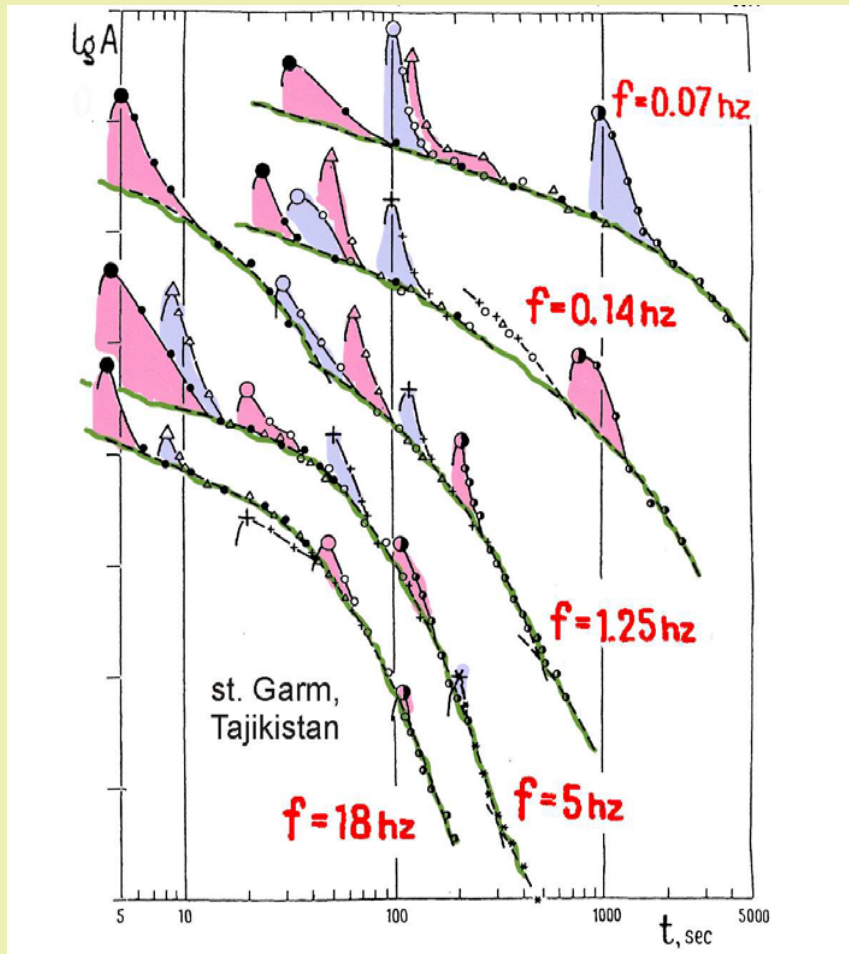
In the 1-12 Hz f range, and for $r=100$ km, MFP estimates are around 100km

Average pulse shapes and their fit by predictions of

- (1) Gaussain-ACF model and
- (2) self-similar inhomogeneity case with $\alpha=3^{2/3}$ (Kolmogorov's spectrum)

The onset-to-peak delay vs frequency relationship indicates $\alpha \approx 3.8-3.9$

Regional envelopes give qualitative understanding of scattering in the Earth(1)



Rautian et al. 1981

(1) Over the entire 20-30 to 400-800 km distance range, the *S*-wave group/pulse is seen *above* coda asymptote.

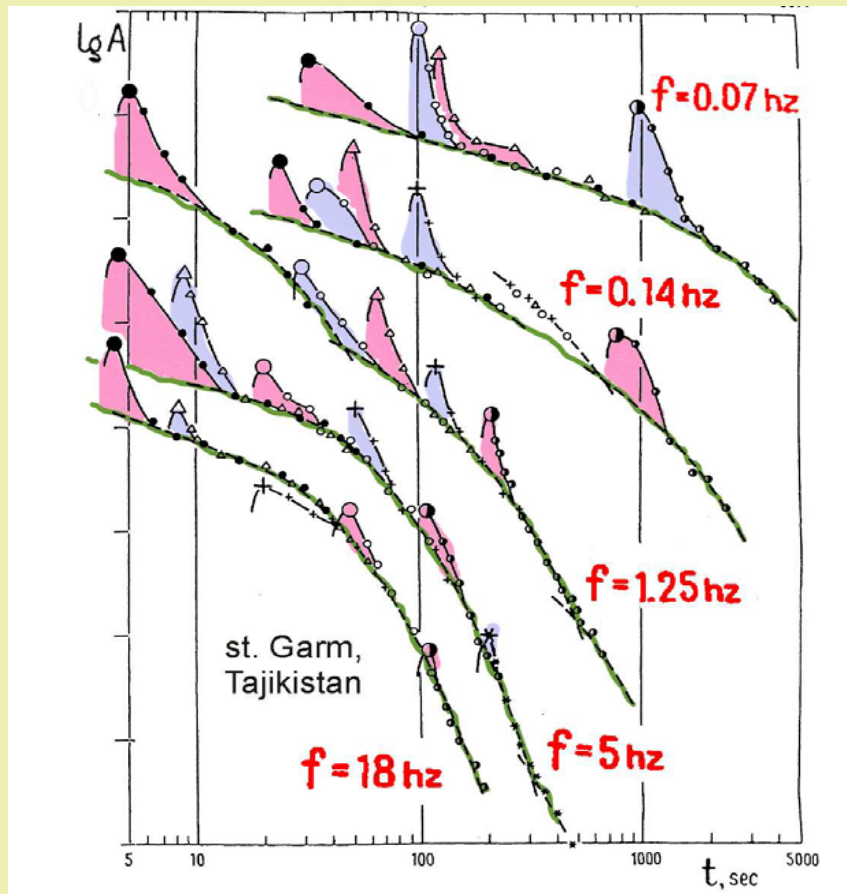
(2) The duration of the pulse is increasing with distance. This pulse broadening is caused by medium, not source, and must be produced by forward-scattering. (Continental Lg is a special case).

(3) Diffusion scattering is not observed. Pulse duration is, roughly, proportional to distance.

(1,2,3) suggests scattering phenomena in general but do not match the picture of scattering in the uniformly scattering medium, (that predicts (a) quadratic trend of duration vs. distance, and (b) fast sinking of a pulse in the diffuse envelope)

All this implies: *ray-average MFP is not constant but rapidly decreases with distance.*

Regional envelopes give qualitative understanding of scattering in the Earth(2)



Rautian et al. 1981

Ray-average MFP is not constant but rapidly decreases with distance. Therefore, in the Earth, for almost any ray and any HF band:

distance r is less than or comparable to ray-average *MFP*

or

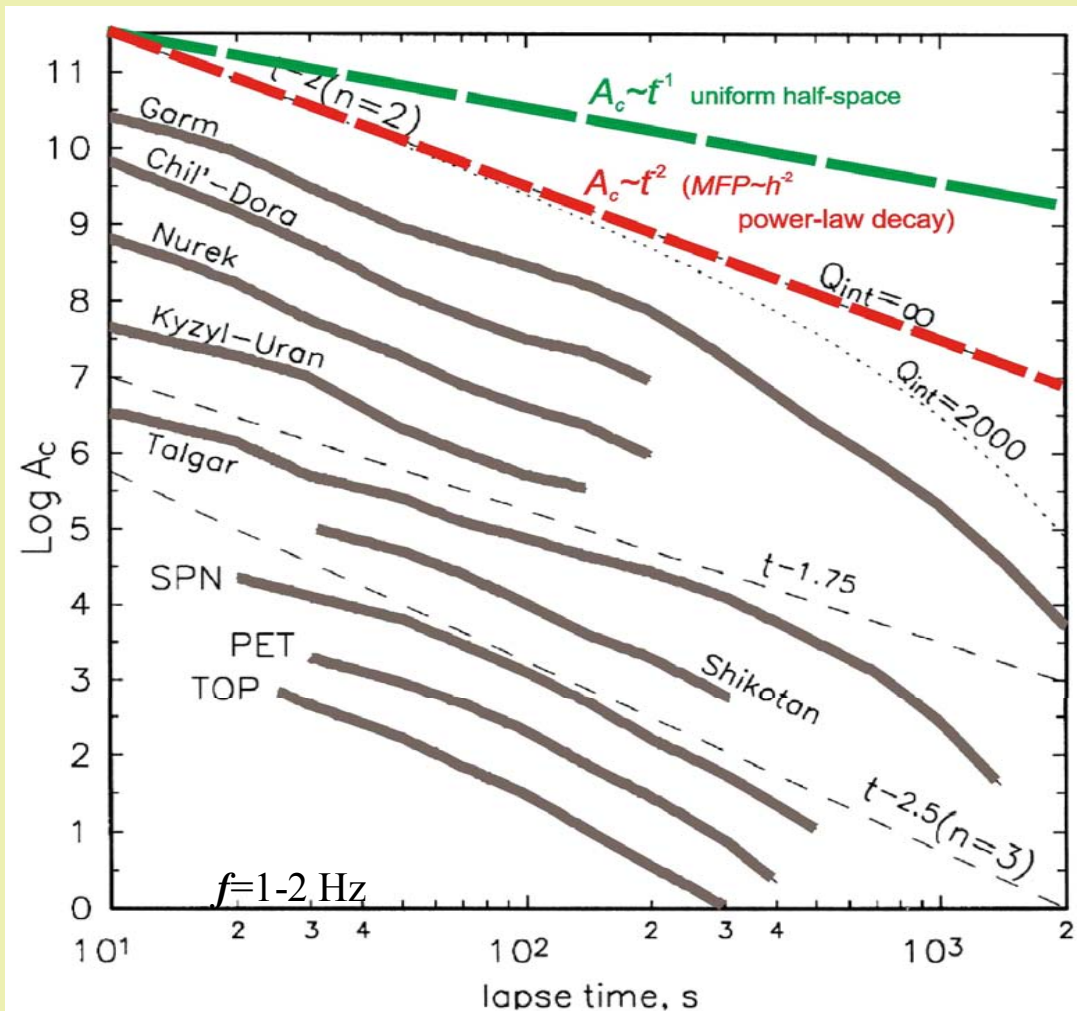
ρ is less than or comparable to 1.0

As rays dive deeper with increasing distance, this means that in the Earth

scattering effects rapidly decay with depth

(follows as well from the existence of impulsive teleseismic P-waves)

Estimating the *transport MFP* vs. *depth* trend from coda shape



Observed coda amplitude over a wide lapse-time range follows neither

t^{-1} (SIS in the uniformly scattering space)

nor

$t^{-1} \exp(-\pi ft / Q_i)$

(same+intrinsic loss labeled “coda Q”).

Instead, a trend like

$t^{-1.75-2.5}$

is seen,

corresponding to SIS in the scattering half-space with very fast depth decay of MFP:

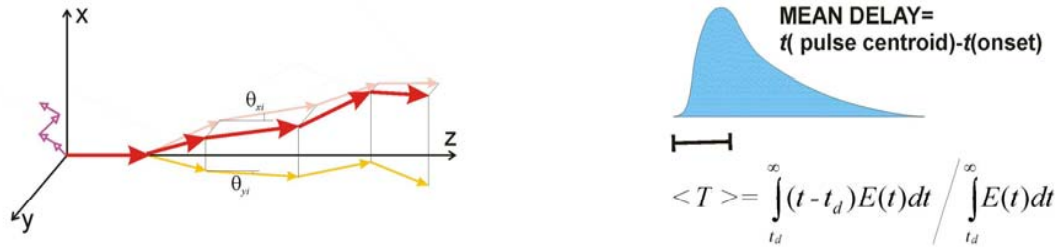
$MFP(h) \sim h^{-1.5-3}$

(adjustment: $\exp(-\pi ft / Q_i)$ with $Q_i = 2000$)

(A traditional coda-Q determination yields a mixture of MFP(h) effect and of intrinsic Q. It can match S-wave Q because a large fraction of S-wave attenuation is caused by radiation loss into deeper weakly scattering layers, thus emulating intrinsic loss in a uniform space.)

Estimating the *transport MFP* vs. *depth* trend from pulse broadening

Basis for inversion:
 mean delay of a pulse = $f(g(\mathbf{r})$ along a ray)



let transport MFP $l=l(\mathbf{r})$, tr. turbidity $g=1/l=g(\mathbf{r})$

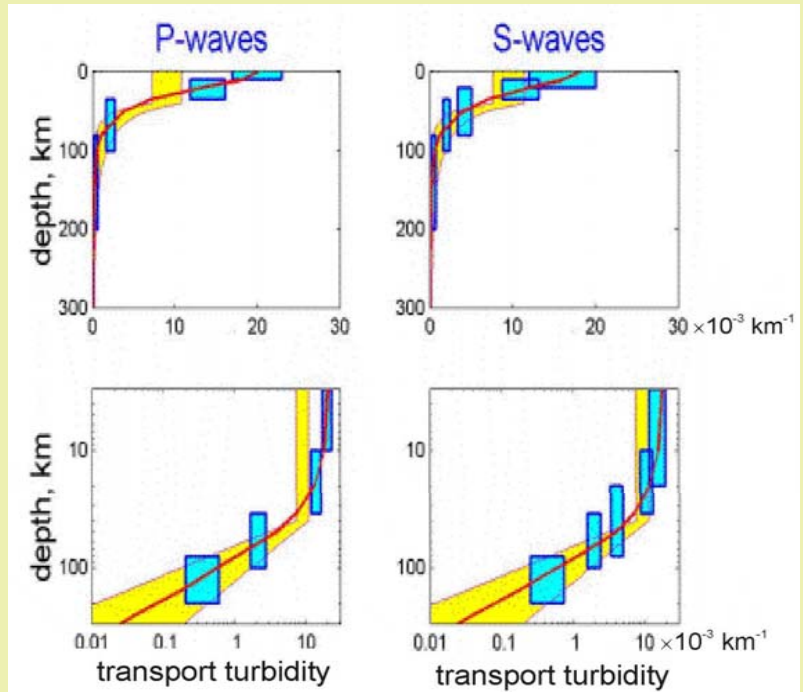
(1) $g(\mathbf{r})=const=g$: $\langle T \rangle = \frac{gr^2}{6c}$ (Williamson 1972)

(2) non-uniform case: $\langle T \rangle = \frac{1}{cS} \int_0^S g(u)(S - u)u du$

where u is the along-ray distance and S is the length of the ray
 (Bocharov 1988)

in practical inversion assuming $\alpha=3.7$
 and thus: onset-to-peak delay = $0.28\langle T \rangle$

inverted vertical profiles $g(h)$
 for P and S waves under Kamchatka
 (based on ~ 2500 onset-to-peak delays,
 from hypocenters at $h=20-300$ km)



colors: different estimates

1. from $h=10-15$ to $h=40-50$ km:
 TMFP $\sim 50-100$ km
2. from $h=60-80$ km down,
 fast decay: TMFP $\sim h^{-2-3}$

VERY IMPORTANT TOPICS NOT COVERED:

1. Conversion scattering: $P \rightarrow S$, $S \rightarrow P$, $S \rightarrow$ surface wave ...)
2. Surface wave (2D) scattering.
3. Inversion of the HF radiation capability function (seismic luminosity) of a finite earthquake source from scattered envelopes

OTHER IMPORTANT TOPICS NOT COVERED :

1. Regional specificity of scattering. Case of Lg
2. Inversion of diffusive envelopes.
3. Synthesis of scattered envelopes.
4. Inversion of observed coda for the relative density of scatterers in 2D or 3D (assuming uniform Q)
5. Inversion of observed coda for the distribution of Q (assuming uniform density of scatterers)
6. Diffraction-based approach (Flatte&Wu 1988)

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The sources of some graphics above:

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M.Fehler. http://www.ees4.lanl.gov/staff/fehler/Fehler_MGSS_Talk_1.pdf

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More complete references see at: <http://www.scatt.geophys.tohoku.ac.jp/>