

The Abdus Salam International Centre for Theoretical Physics







SMR.1676 - 47

## 8th Workshop on Non-Linear Dynamics and Earthquake Prediction

3 - 15 October, 2005

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Unified Laws in Seismicity (copies of transparencies)

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These are preliminary lecture notes, intended only for distribution to participants

Unified Scaling Laws in Seismicity G. Molchan IIEP (Moscow), ICTP (Trieste) 03-15 October 2005

Outline

motivation

• examples

• existence & shape

• multifractality

• optimal spatial scaling

Scale free distributions • power laws: GR, Omori • unified scaling laws (Baketal, 2002) 3 - statistics  $\Delta$  - scale (space or time or both)  $\exists a_{\Delta} : a_{\Delta} \xi_{\Delta}$  is independent of  $\Delta$ p.d.f. of  $\mathcal{Z}_{\Delta} = a_{\Delta}f(a_{\Delta}x)$ where  $f(x) = p.d.f.of \eta = a_0 \xi_0$  $a_{\Delta} = c/E \sharp_{\Delta}$ 

Reasons for consideraton:

- · evidence of complex spatio-temporal behaviour of seismic regime
- a tool for renormalization of precursors
   de-clustering of EQ catalogs
- · understanding of limits in predictability

Problems

- · existence of unified scaling laws

- shape
  real scale range
  new information contained in a unified law



Fractality of epicenters λ(dg) - rate of events of M>me in an area dg **図** - 入(L×L)>0  $\times \{ \square \} = n_{\perp}$ Definition : d, is the Box-dimention of support of measure  $\lambda(dg)$  $\log n_L \simeq d_f \log Lo/L$  , Leel  $\frac{1}{2} = \begin{bmatrix} 1.8 & m>2 & L=10-100 \\ m>3.5 \\ 1 & 10 \\ 1 &$ L=100-1000 Other estimations: 1.6 (Corral) 1.2 (Kagan, Kossobokov

Consequence (Keilig-Borok et al, 1989)

 $\langle \lambda(L \times L) \rangle \approx \text{const } 10^{-8m} (L/L_0)^{d_f}$   $\uparrow$   $\lambda > 0$ 

Problems: range of L df is not unique; df = df(m)-? Correlation of events

Baiesi & Paczuski , 2004



FIG. 2: The probability distribution of the correlation,  $c_{,}$  between all earthquake pairs in the data base, with  $m_{<} = 3$ , using both the 2D metric and the 3D one. They are scale free distributions over many orders of magnitude. The threshold  $c_{<} = 10^{4}$  where correlations are considered significant and links are made is indicated in the figure. Note that, with that threshold, most links are eliminated from the network, giving a reduced data set to examine seismic properties.

[9] meter , [+] sec , 1gA = -4 (1+dimg)



Scaled distributions of rates,  $\lambda(L \times L)/a_{L}$ , for several At, L, and me. dy=1.6 , b=0.95 a1 = 10- 6m (L/Lo <u>AL = 8-140 km</u> At = Iday - gyears

Waiting-time distribution: unified law by Baketal (2002)



d1 - ?

Statistics T2: time betweenLxLconsecutive EQsof M>mc in a random cell LxL

Probicell LxL } =  $\lambda(LxL)/\lambda(G)$ 

Normalization: East\_ = 1

 $a_{L} = \langle \lambda(L \times L) \rangle = c \left( 0^{-bm} \left( \frac{L}{L_{0}} \right)^{d_{f}} \right)^{d_{f}}$ 



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Scaled distributions of 
$$T_{L}$$
 (Corral, Phys. Rev E 68,2003)  
 $a_{f} = 1.6$ ,  $b = 0.95$ ,  $T > 38s$   
 $\Delta L = 15 - 250 \text{ km}$   
 $\Delta m_{c} = 2 - 4$ 

5° Unified Scaling Law by Bak Bak's philosophy applied to EQs: Don't care about • the tectonic environment (a subdivision of seismogenic region is independent of it)

- after-, for- or main-shocks
- (there is nothing in the seismograms that differentiates these events)
- · space and temporal heterogeneity

Waiting-time distribution: Universal scaling law by Corral statistics T<sub>G</sub>: time between consecutive EQs of M>m in a region G G

Normalization :  $E a_L T_L = 1$  $a_L \not \sim 10^{-8m} (L/L_0)^{d_f}$ 



FIG. 1. Local distributions of interoccurrence times for several stationary periods and different regions, after scaling by the rate.



Corral: arxiv: cond-mat/0310407 17 Oct 2003

Theoretical Results (Molchan, 2005)

1° The Shape of Corral's universal scaling law

If there are two areas with independent seismicity then a universal p.d.f. of ZL, ETL=1, is exponential

more exactly:

$$G_{1} \vee G_{2} = G_{0}$$

$$N_{i}(dt) - stationary sequence of events i G_{i}$$

$$N_{0} = N_{1} + N_{2}$$

$$N_{i} \text{ and } N_{2} \text{ are stochastically independent}(!)$$

$$+ the rescaled T-distribution for G_{i}, i=0, 1, 2$$

$$are the same with p.d.f. f(x)$$

$$+ f(x) < cx^{\Theta}, 0 < \theta < 1, x < 1$$

$$then \qquad f(x) = exp(-x)$$

$$(onclusion: observed f(x) \neq exp(-x))$$

Corral's universal scaling law can't exist in the whole range of time

## 2° Main peculiarities of the rescaled TG distribution

In the general Poissonian cluster model with Omori (aw  $\Lambda(t)$  1  $\infty$ ,  $t \ll 1$ p.d.f. of  $a_G \tau_G(t) \propto \begin{pmatrix} exp(-t/a) \ t \gg 1 \\ \Lambda(t) \end{pmatrix} = \frac{t}{t}$ 

where a is fraction of main events among all seismic events

• (\*) is supported by observations

• What other information can be extracted from the "universal" law by Corral !



Spatial scaling of seismicity zate  

$$\lambda(L \times L)$$
 is rate of  $\{M > m_e\}$  in cell  $L \times L$   
Scaling:  
 $\lambda(L \times L) \propto L^{+d}$   
 $T(L \times L) = O(1/\lambda(L \times L)) \propto L^{-d}$   
Practice:  $d = d_0$  (Box dimension)  
 $d_2$  (correlation dim)  
 $d_p$  (generalized dim  
by Grassberger- Procaccia)  
Problem: which d is better?  
Example(Pisarenko 1996)  
If  $\lambda(dg)$  is a sample of a random Levy measure  
of index  $0 < d < 1$  then  
 $P(-\lambda(L \times L)/L^{2/M} < X)$   
is independent of  $L_{-1}$  i.e.  $d = 2/d > 2$   
Here  $E \lambda(L \times L) = \infty$ ?  
 $d_p \neq const \iff \lambda(dg) - multifractal measure$   
?

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H-dim (Hausdorff dimension) is more strict analogue of B-dim

Singularity of typed Hoelder's exponents 2° gELxL  $\log \lambda(L \times L) \simeq \lambda \log L$  ,  $L \to 0$  . Z d=1 **√**=2 Sa={g: g is of type a y 3° 2(dg) is multifractal if  $S = \{ support of \lambda(dg) \} = U S_d$ ,  $s_d \neq \emptyset$ Multifractal Spectrum: (d, f(d))  $f(d) = \dim S_d$ 



trivial mixture of mono fractals (no similarity) f(d):d , d= do, 1, 2

1° 
$$T(q)$$
 function  
Renyi function  
 $R_{L}(q) = \sum_{\lambda>0} \left[ \frac{\lambda(L\times L)}{\lambda(G)} \right]^{q}$  191 <--  
 $\log R_{L}(q) = T(q) \log L$ , L=0  
Multifractal formalism:  
 $T(q) = \min(qd - f(d)) := \chi f$   
If  $T(q)$  is strictly concave and smooth  
then  
 $f(d) = \min(qd - T(q)) := \chi T$   
 $(\chi - the Legendre transformation)$   
•  $\min \rightarrow d = \tilde{T}(q) \rightarrow d$  is unique (monofractal)  
 $q = \tilde{T}(q) = d(q-1)$  is linear  
• Box-counting method give H-dim(Su)!

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$$d_{q} = \frac{T(q)}{q-1} = \frac{T(q)-T(q)}{q-1} = \frac{T(q^{*})}{Hoe!der's exponent}$$

$$d_0 = -T(0) - Box-dim$$
  
 $d_1 = \dot{T}(1) - the information dimension$   
 $d_2 = T(2) - the correlation dimension$ 

1-5

Typical T,f functions





t(1) exists ⇒ do, d1 are dimensions of suppx(dg) other dp are Hoelder's exponents

Problems :

- multifractality of real seismicity, i.e.  $T(q) \neq d_0(q-1)$  (evidence)
- optimal scaling of X(1×1) and T(1×1)
   for random cell 1×1 under multifractal
   Condition

Random cell LxL:



R - L×L  $Pr(\blacksquare) = C_p[\lambda(L\times L)]^P, P = 0$  $1/c_{p} = \sum [\lambda(1 \times L)]^{p} = R_{1}(p) / [\lambda(G)]^{p}$ 

P=1 - the Bak's Case

P=0 - the case of generalized G-R law $<\lambda(L \times L) > = \frac{1}{n_L} \sum_{\lambda > 0} (L \times L)$  $and of distribution of {\lambda(L \times L)}$ 



California: seismic events with M≥3 (·); grid of spacing L=100 km (°) centered at (rombus); the main (G) and alternative (G<sub>1</sub>) rectangular seismic region for dimension computations



Tau-function for M≥2 events in region G; it is based on the interval of scales  $\Delta L$ =(10, 100) km. The broken straight line of two segments

is  $\tau(q)$  for a bi-fractal

, with the observed dimensions  $d_0$  and  $d_2$ .

 $d_{o=1.8\pm0.1}$ ,  $d_{1} = 1.35\pm.05$   $d_{2} = 1.1\pm.05$ 

Evaluation of 
$$T(q)$$
  
log  $R_L(q) = T(q) \log L$ 

Problem: find the range of scale (L, 2\*) with stable slope in (Log RL(9), logL)

Idea: isolated points in support of λ(dg) do not affect its multifractal spectrum Use cell L×L with ¥ EQs > K=0,1,2...

• n2 7,100 ⇒ L<2"



Data for estimating the box dimension  $d_0 = -\tau(0)$ based on M  $\ge$  2 events (a) and M  $\ge$  3 events (b).

The vertical axis snows the number of L×L cells, n(L|k), that have numbers of events greater than k, k = 0, 1, 2, 3, 4. The dashed line shows both the slope  $\tau(0)$  and the interval of scales  $\Delta L$  for estimation of d<sub>0</sub> by least squares using n(L|0).



Data for estimating the information dimension  $\sigma_1$  from M≥2 (a) and M≥3 (b) events

The vertical axis shows the derivatives d/dq  $R_{L}(q=1|k)$  for k=0, 1, 2, 3, 4 where  $R_{L}$  is the modified Renyi function.

The dashed line shows both the slope  $d_1$  and the interval of scales  $\Delta L$  for estimation of slope by least squeres.

## Weighed means of $\lambda(L\times L)$ and $\overline{\delta}(L\times L)$ : Scaling $W^{(p)}(L\times L) = [\lambda(L\times L)]^{p} c_{p} - weight of sell L\times L$ 1° $(L\times L)^{p} = \sum_{\lambda>0} \lambda(L\times L) \cdot W^{(p)}(L\times L) \sim L^{d}$ $d = \overline{\delta}(p+1) - \overline{\delta}(p) = \begin{cases} d_{0} & p=0 \\ d_{2} & p=1 \end{cases}$ (alifornia) $d = \overline{\delta}(p+1) - \overline{\delta}(p) = \begin{cases} d_{0} & p=0 \\ d_{2} & p=1 \end{cases}$

$$2^{\circ} < \overline{\tau}(L \times L) >_{p} = \overline{L} \times^{\prime}(L \times L) \times^{(p)}(L \times L) \sim L^{\widetilde{d}}$$
  
$$\overline{d} = |\overline{\tau}(P-1) - \overline{\tau}(P)| = \begin{cases} |\overline{\tau}(-1)| - d_{\circ} & p = 0 \\ d_{\circ} & p = 1 \end{cases} = 1.8$$
  
$$3^{\circ} \quad d[\langle \lambda >_{p}] \leq \overline{\tau}(P+0) \leq \overline{\tau}(P-0) \leq d[\langle \overline{\tau}_{L} >_{p}]$$

Hint  

$$\langle \lambda^{\varepsilon}(L + L) \rangle_{p} = \lambda^{\varepsilon}(G) R_{L}(P + \varepsilon) / R_{L}(P)$$
  
 $\sim L^{T}(P + \varepsilon) - T(P)$ 

of  $T(L \times L)$  and  $\lambda(L \times L)$ 

Statistics Te: time between consecutive EBs of M>me in Wp-random cell LKL

 $\mathcal{E}_{L} = \mathcal{L}^{d} \overline{\mathcal{I}}_{L} , d^{-?}$ 

Statistics XL: rate of Wp-random cell LX2

 $\mathcal{B}_{L} = \mathcal{L}^{-d} \lambda_{L}$ ,  $d^{-?}$   $\|d \text{ is inadmissible if}$   $\mathcal{B}_{L_{i}}, \mathcal{L}_{i} \rightarrow 0$  can have limit (in distribution)  $\|d f + dype \quad 0 \text{ or } \infty$ Problem 'find all admissible exponents d

are given by unique interval [t(p+0),t(p-0)] i.e. d=t(p) if t exist

d = 1 2(0)	P=0	California 2
( d.	P=1	1.35
l Ż(2)	> p=2	v



Distribution functions for lg  $\xi_{L}$  corresponding to the scales L=10, 20, 40, 80 and 100 km and to the scaling index c = 1.1, 1.35, 1.6, 1.8 and 2.0. The segment ( $\alpha$ ,  $\beta$ ) has the slope (-1), its length provides information on the scatter of the distributions of lg  $\xi_{L}$  at a fixed c.

The case p=0: each L×L cell enters in the distribution with the same weight



Distribution functions for lg  $\xi_{\rm I}$  corresponding to the scales L=10,20,40,80 and 100 km and to the scaling index c=1.1,1.2,1.35,1.6,1.8 and 2.0. The segment ( $\alpha$ , $\beta$ ) has the slope (-1), its length provides information on the scatter of the distributions of lg  $\xi_{\rm L}$ at a fixed c. The case p=1: each L×L cells enters in the distribution with a weight proportional to the seismicity rate in the cell.

Conclusion

 Unified scaling laws by Corral and by Baketal Can't exist in strict sence
 We find a suitable exponents d in the relations T(L×L) αL<sup>d</sup> λ(L×L) α L<sup>d</sup>
 assuming multifractality of rate λ(dg); d is not unique and depends of a problem, i.e. d is a parameter in appications
 We confirm the multifractal property of rate in the range of scale : 10-100 km in California We do it twice : estimating T(g) and using prediction of exponents in scaling of Z and λ<sub>L</sub>