



The Abdus Salam
International Centre for Theoretical Physics



H4.SMR/1775-44

**"8th Workshop on Three-Dimensional Modelling of
Seismic Waves Generation, Propagation and their Inversion"**

25 September - 7 October 2006

Geodynamics of Mediterranean

C. Doglioni

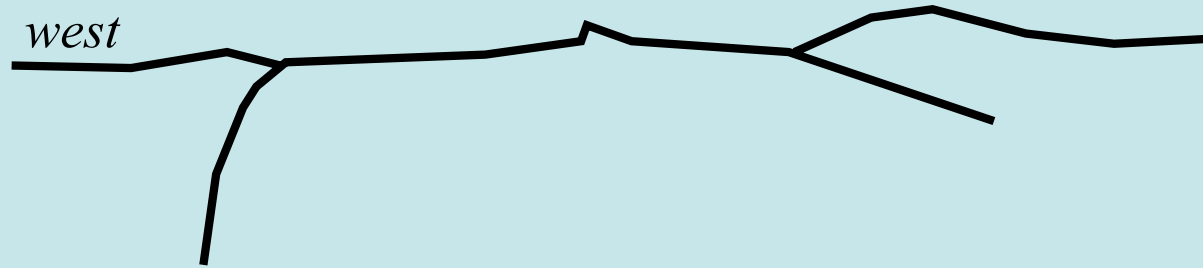
**Universita' degli Studi di Roma
Roma**

Global Tectonic Asymmetry & Applications to the Mediterranean



Imagination rules the world
Napoleon

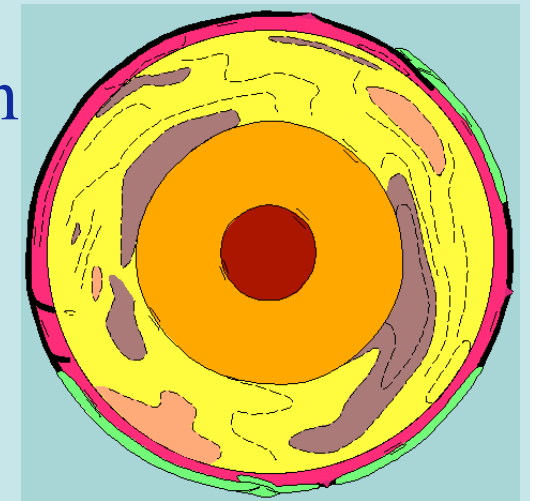
1) Orogens and Rifts show an “E-W” global asymmetry



2) The lithosphere moves along a westerly polarized flow

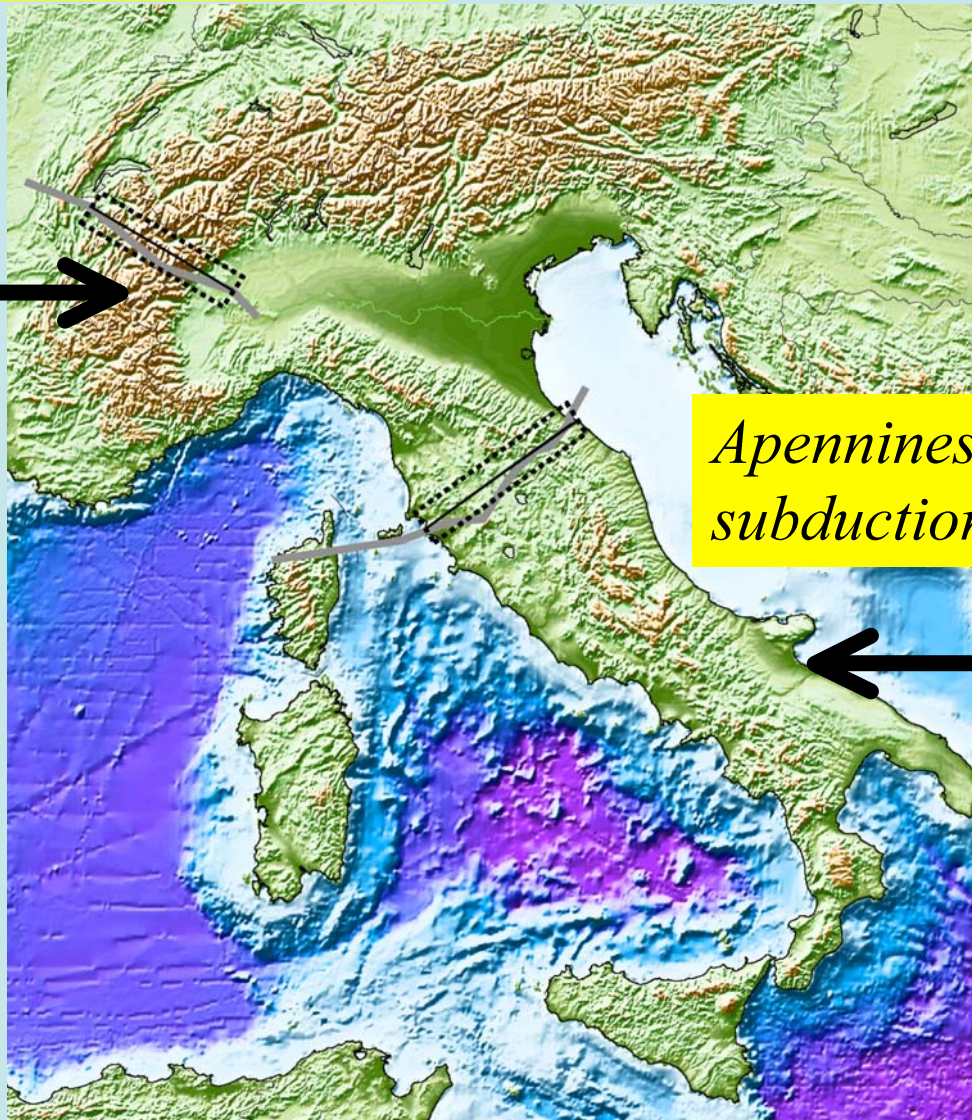


3) Plate tectonics is tuned by Earth's rotation

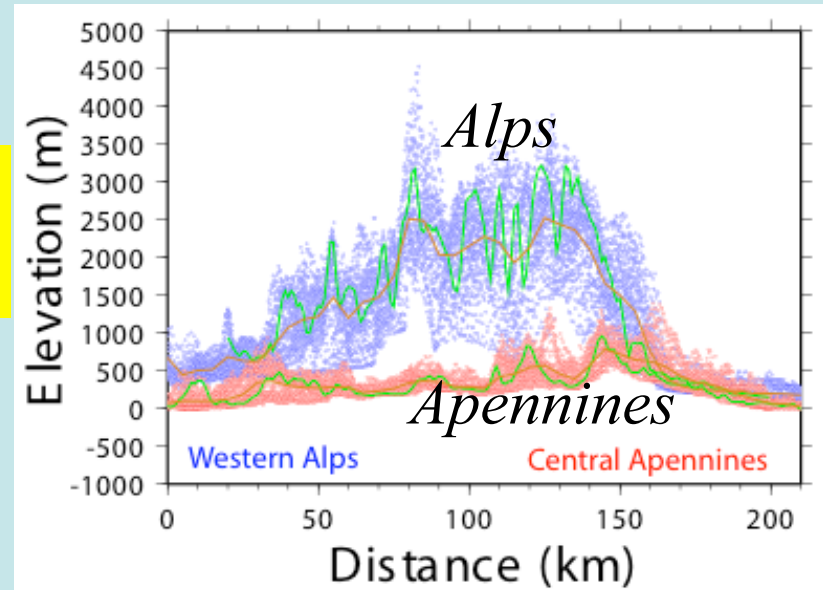


Alpine subduction

TOPOGRAPHY

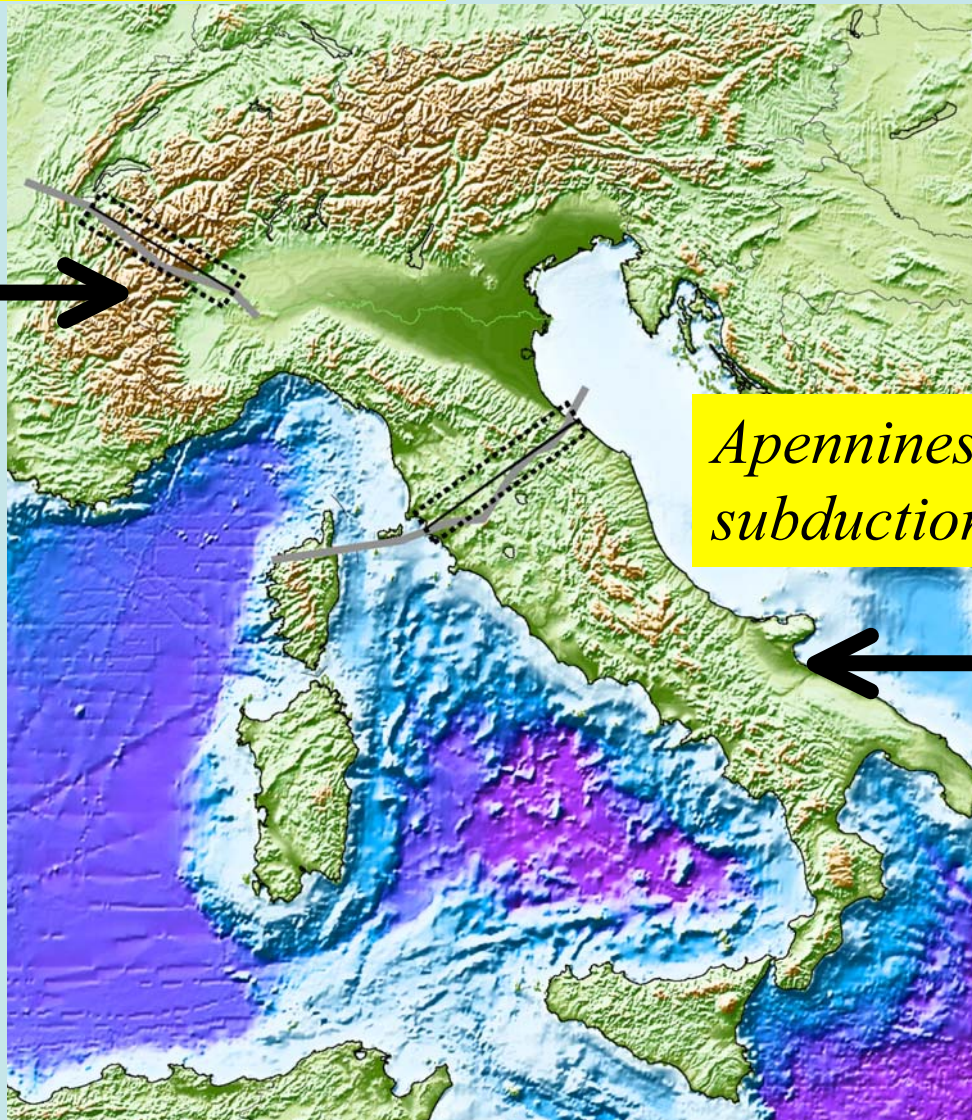


Apennines subduction

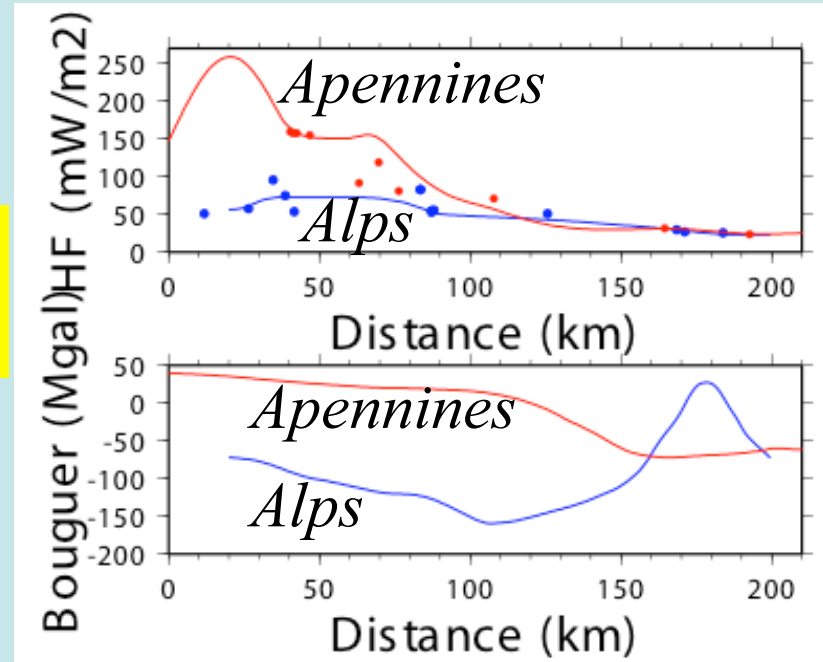


BOUGER ANOMALY & HEAT FLOW

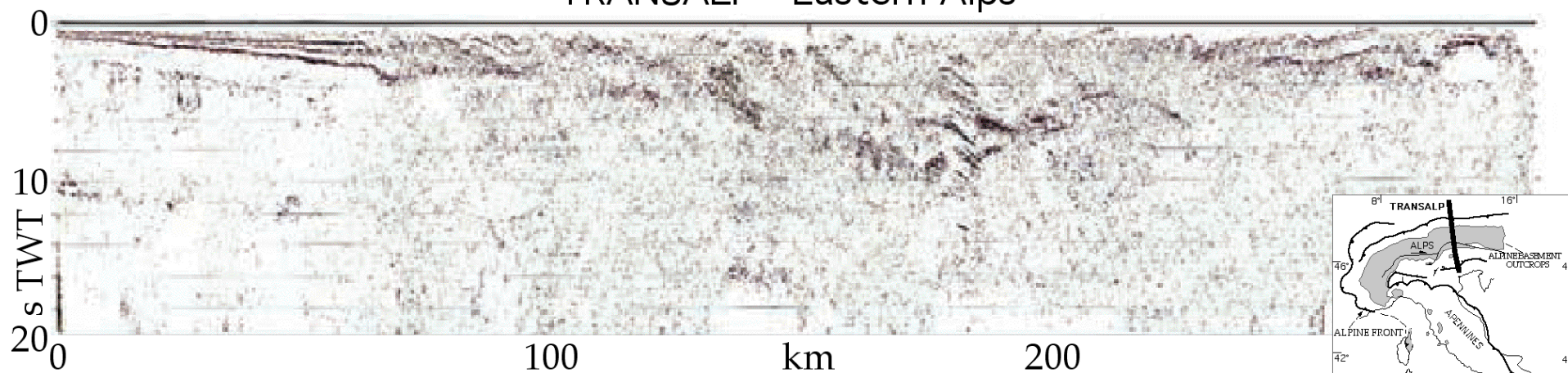
Alpine subduction



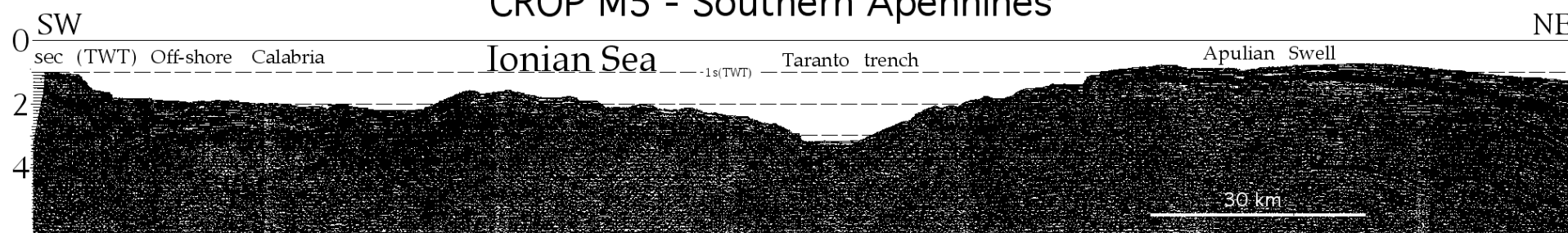
Apennines subduction

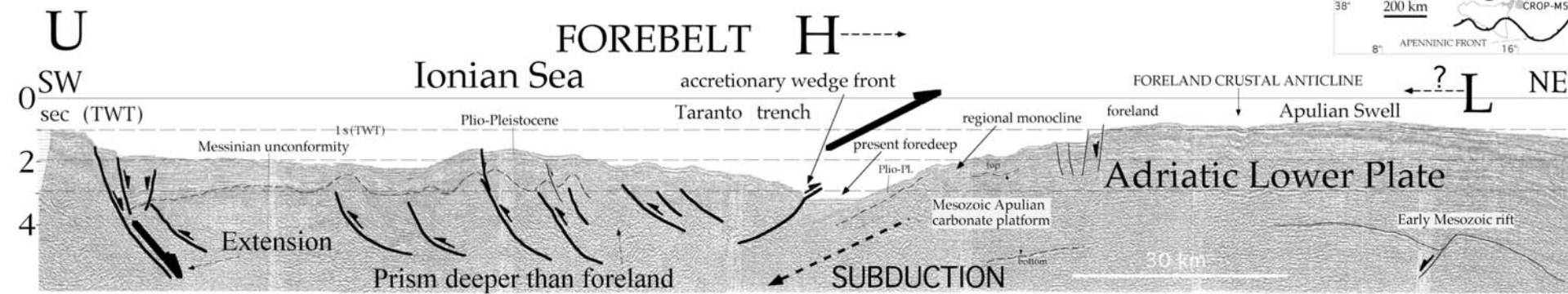
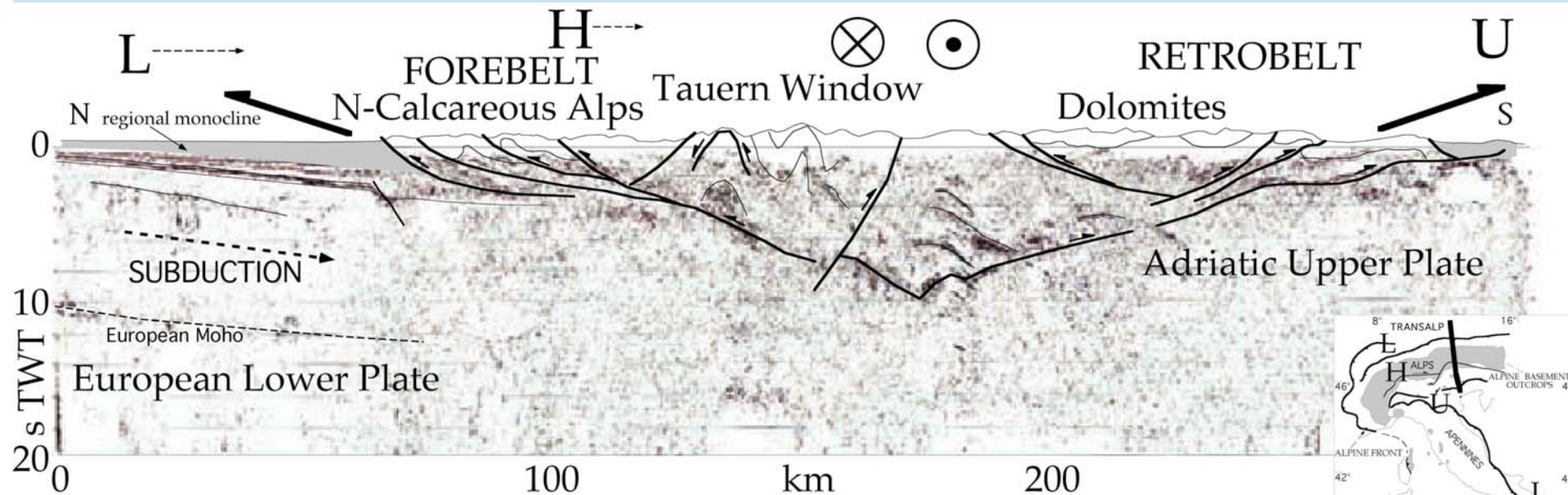


TRANSALP - Eastern Alps

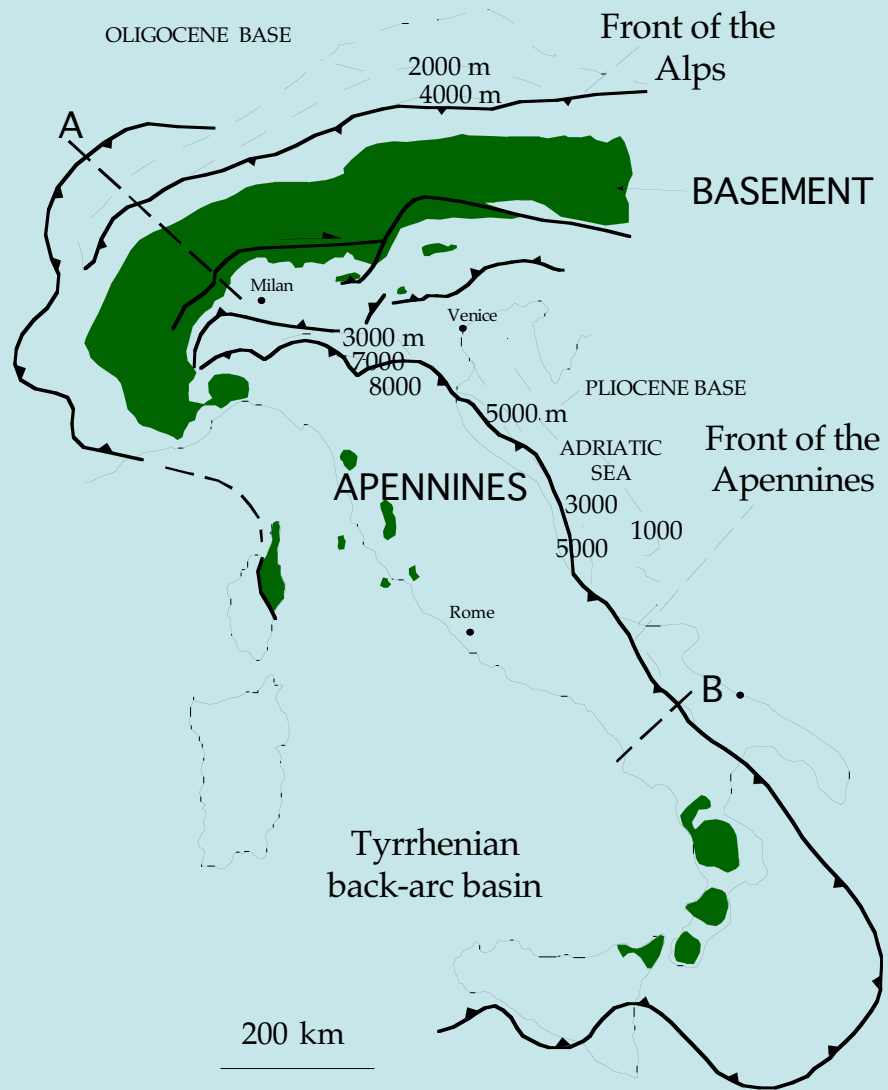


CROP M5 - Southern Apennines

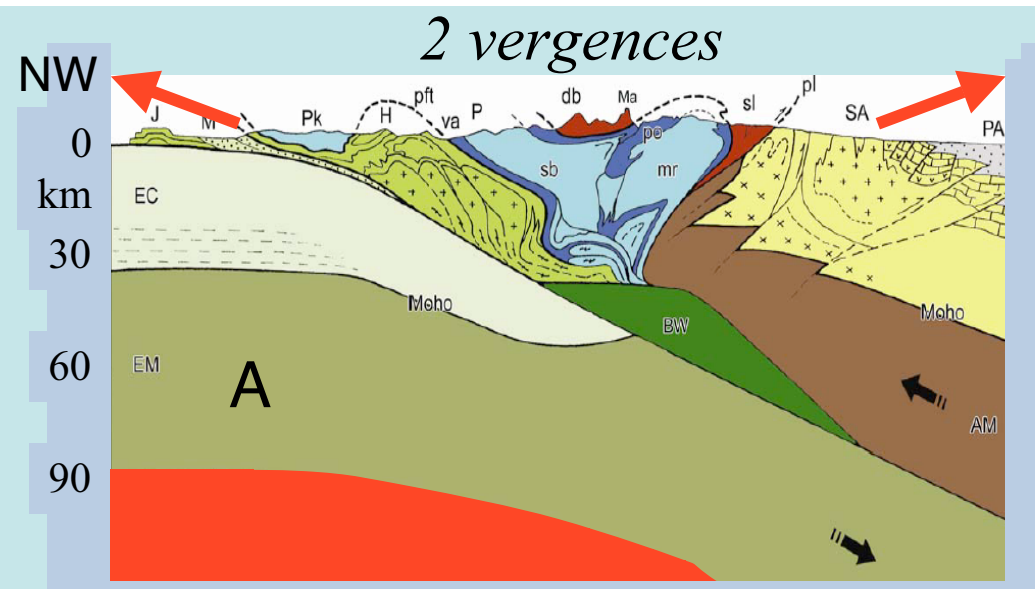
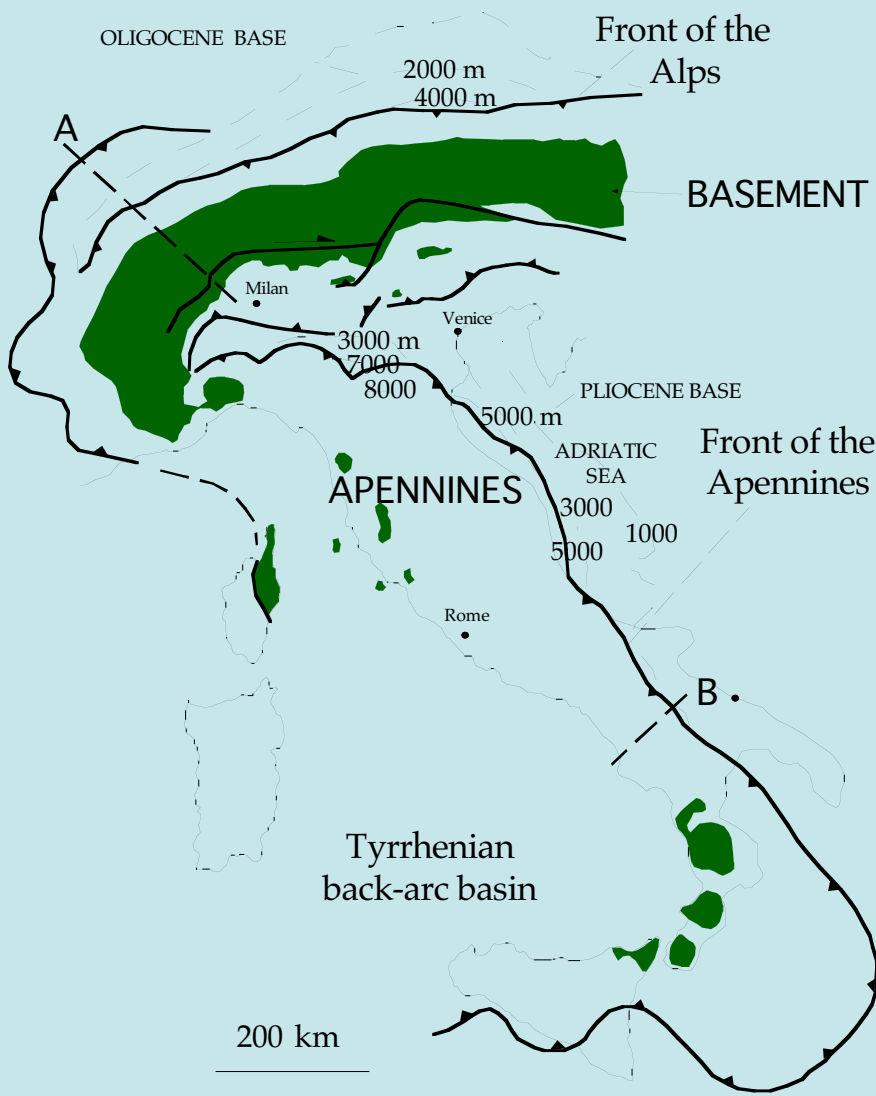




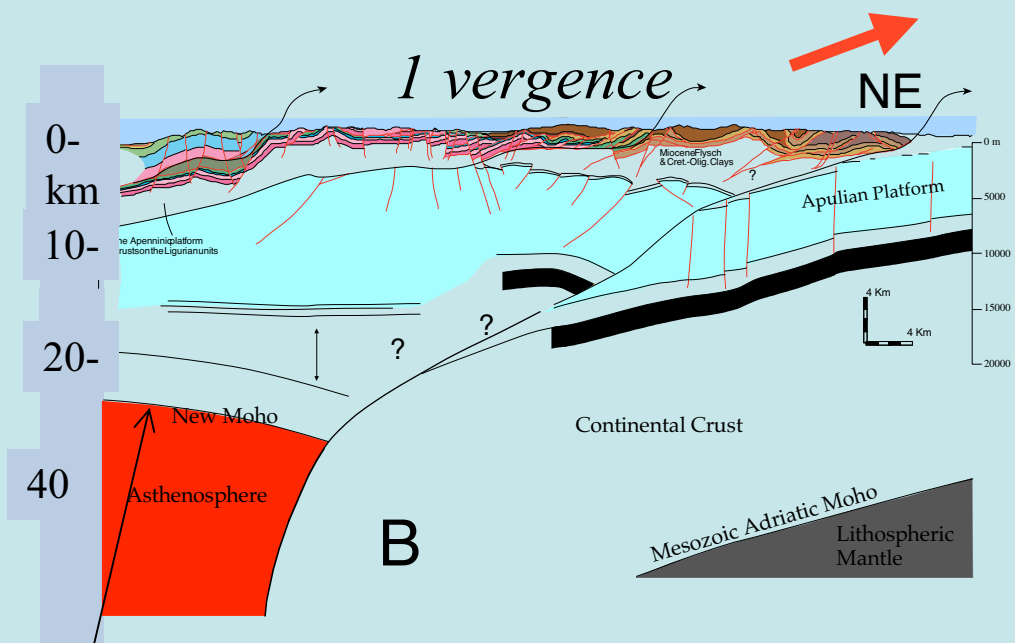
ROCKS



VERGENCES & MOHO



Dal Piaz, 1997



shallow asthenosphere & new Moho



ALPS

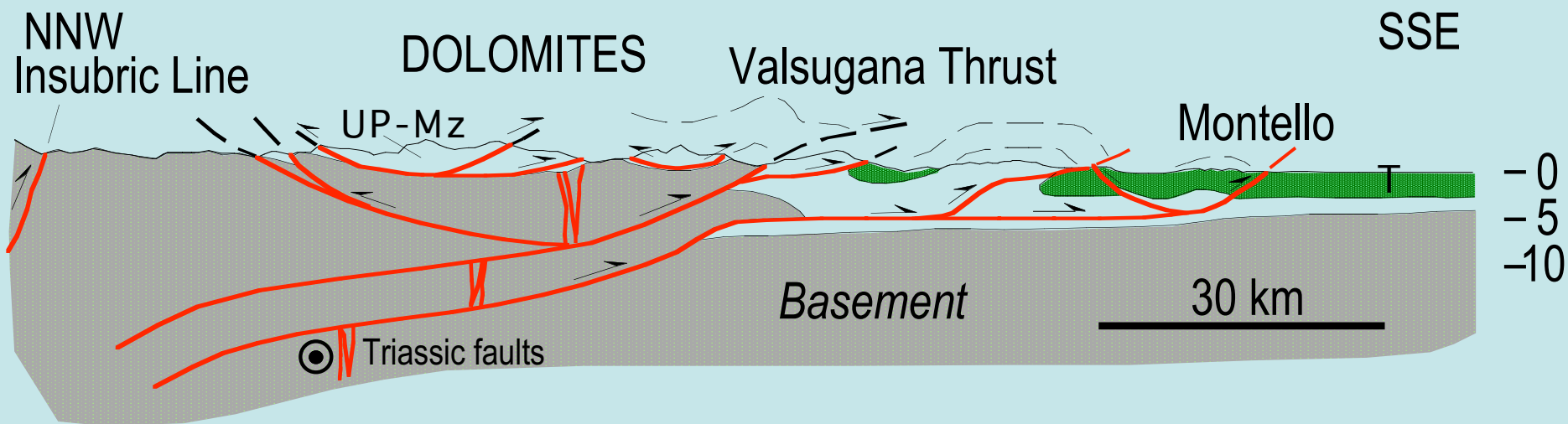
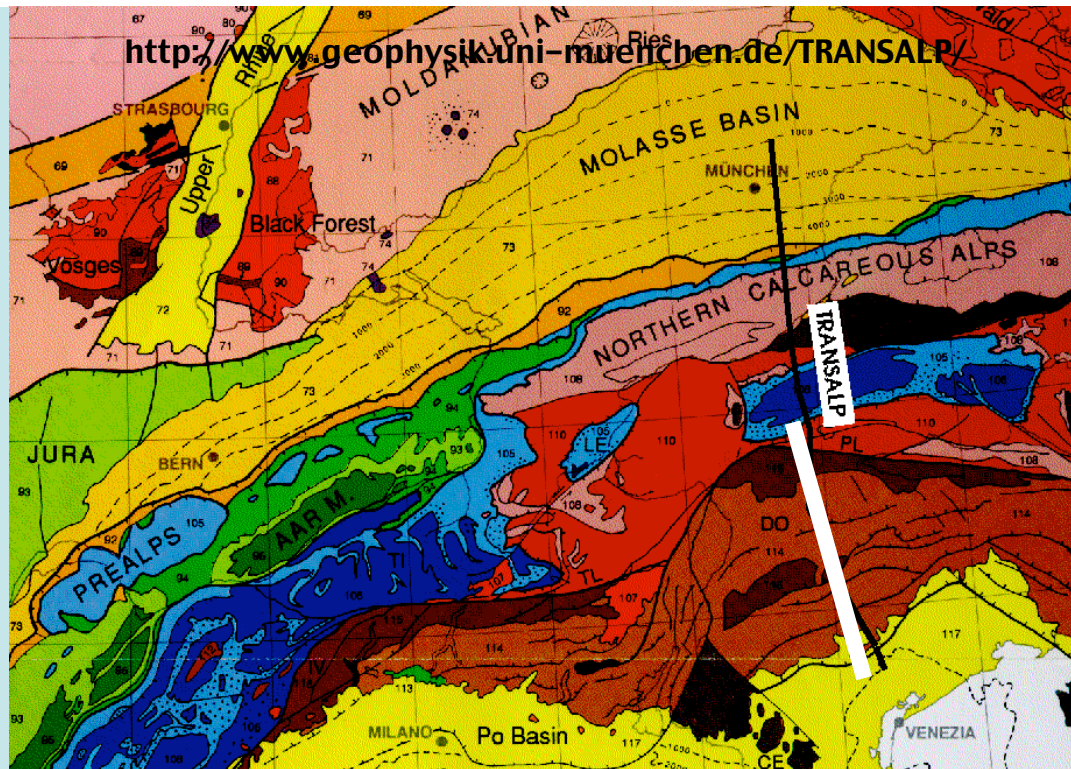
Blue schists ophiolites

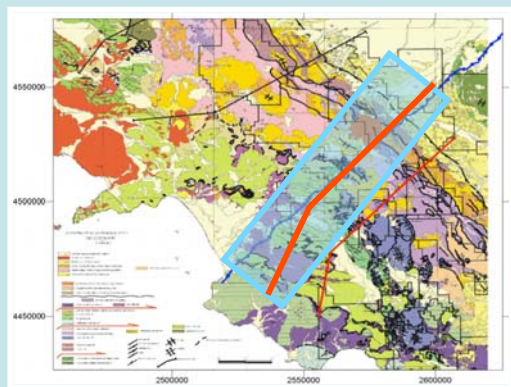
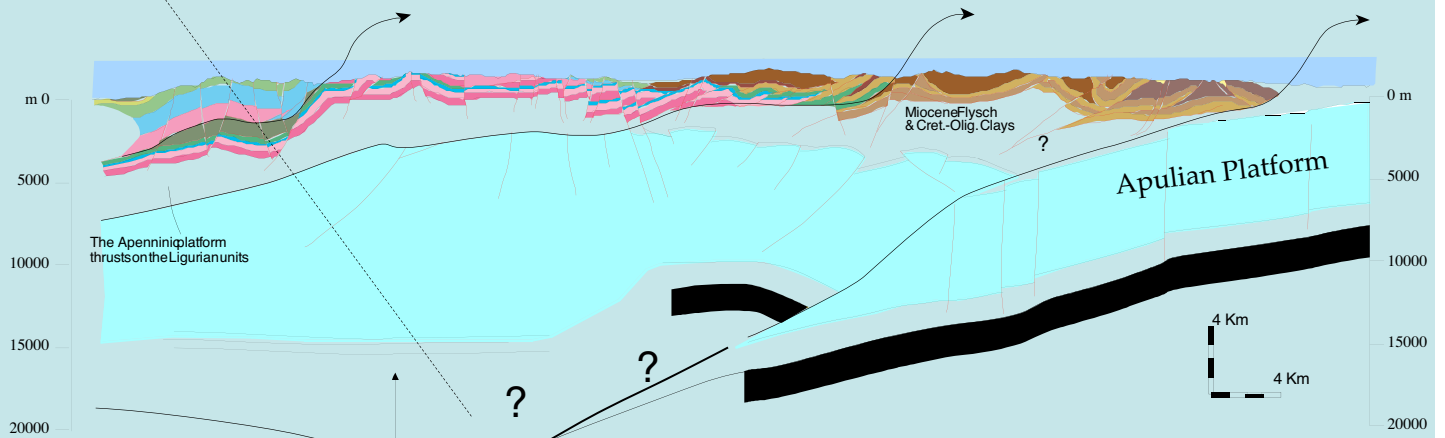
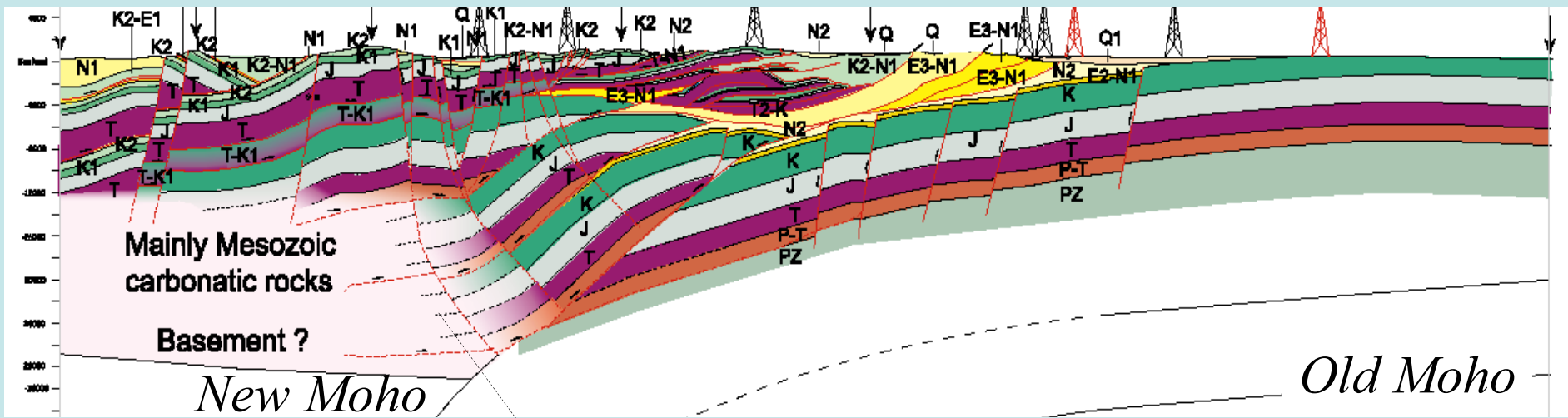
Age 160 Ma - HP 50 Ma

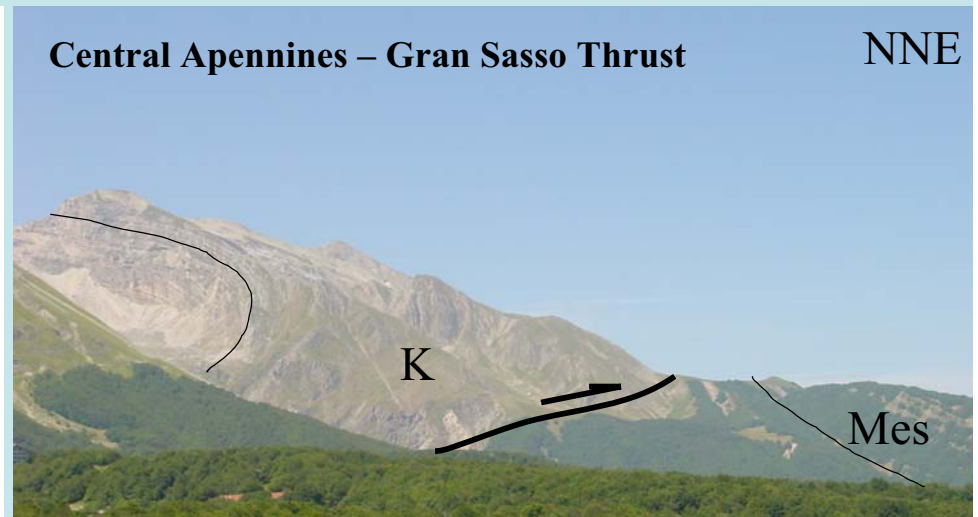
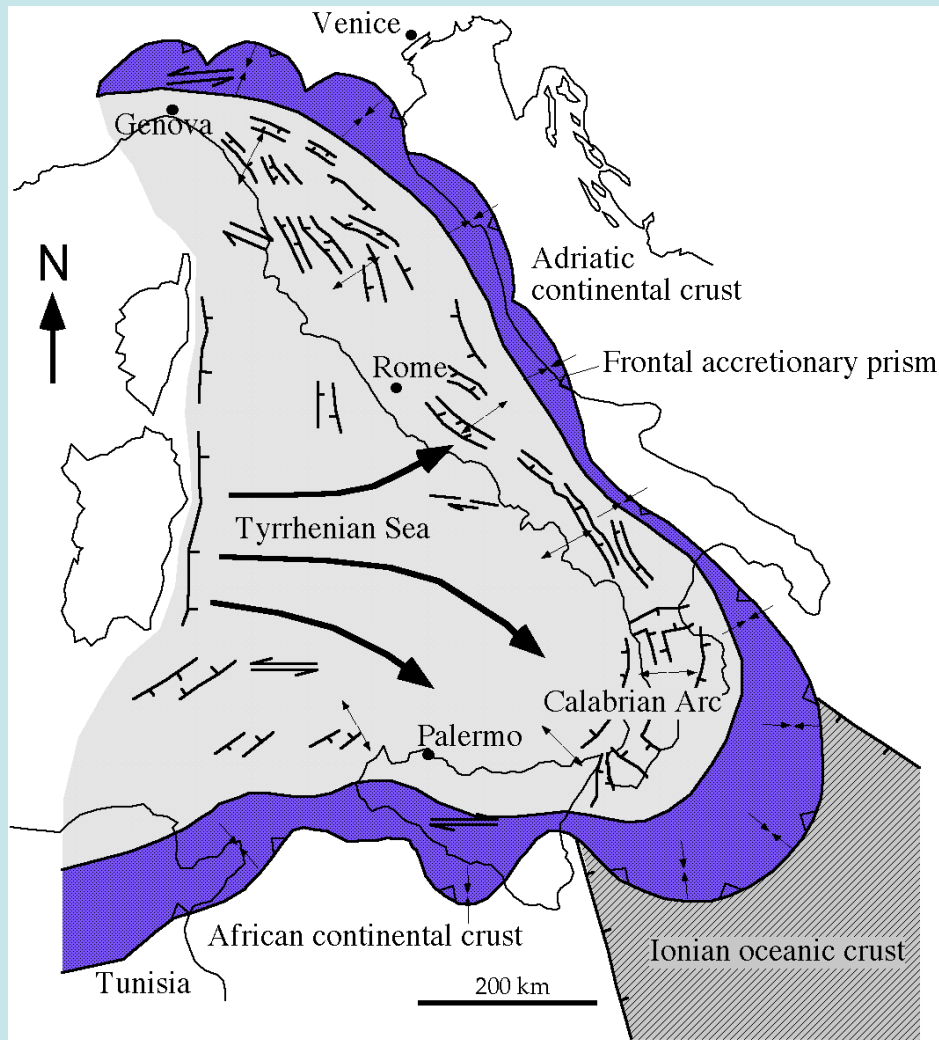
APENNINES

“unmetamorphosed” ophiolites

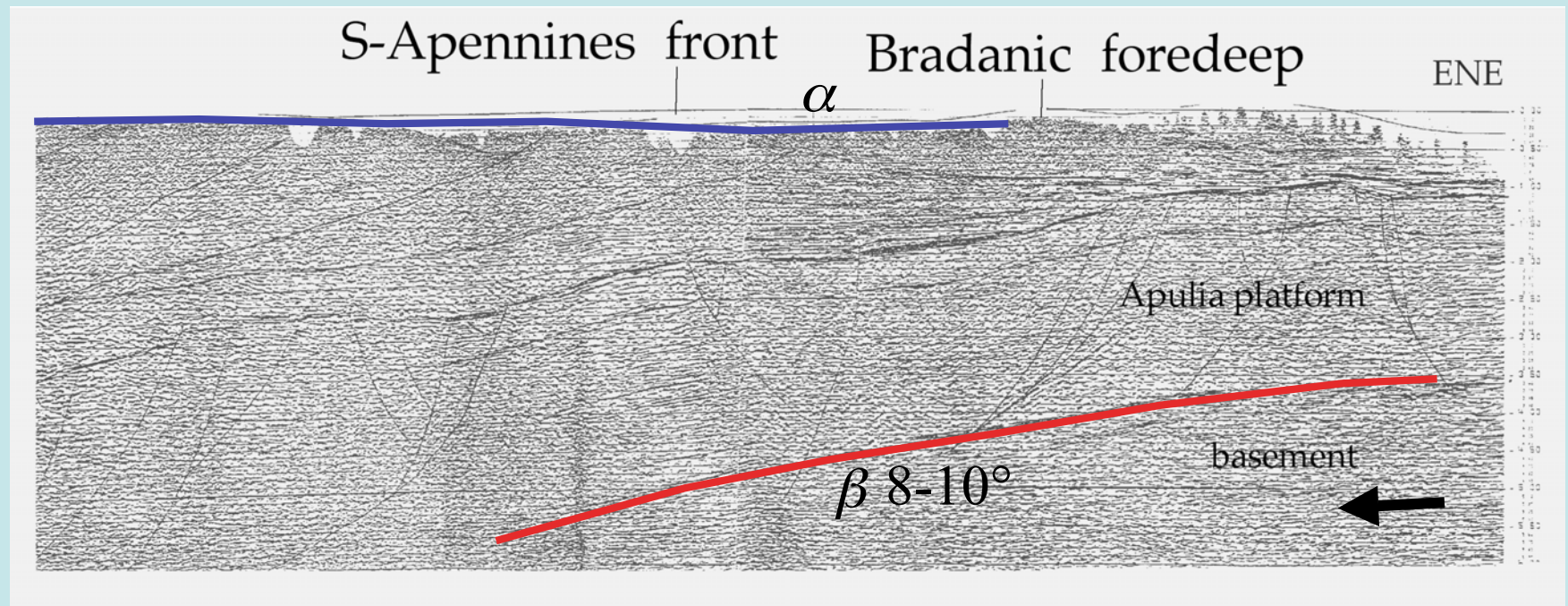
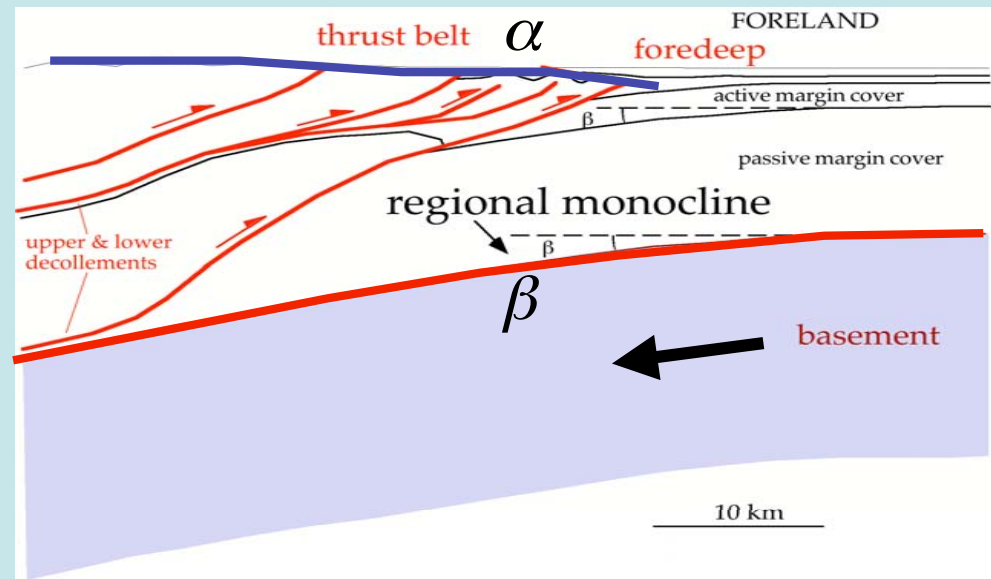
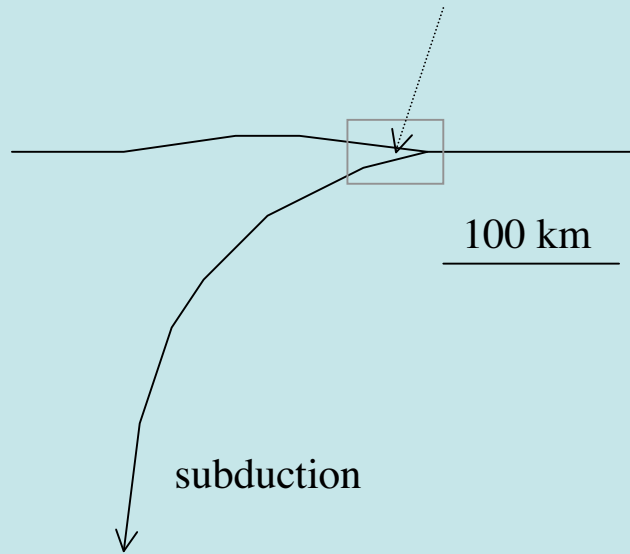




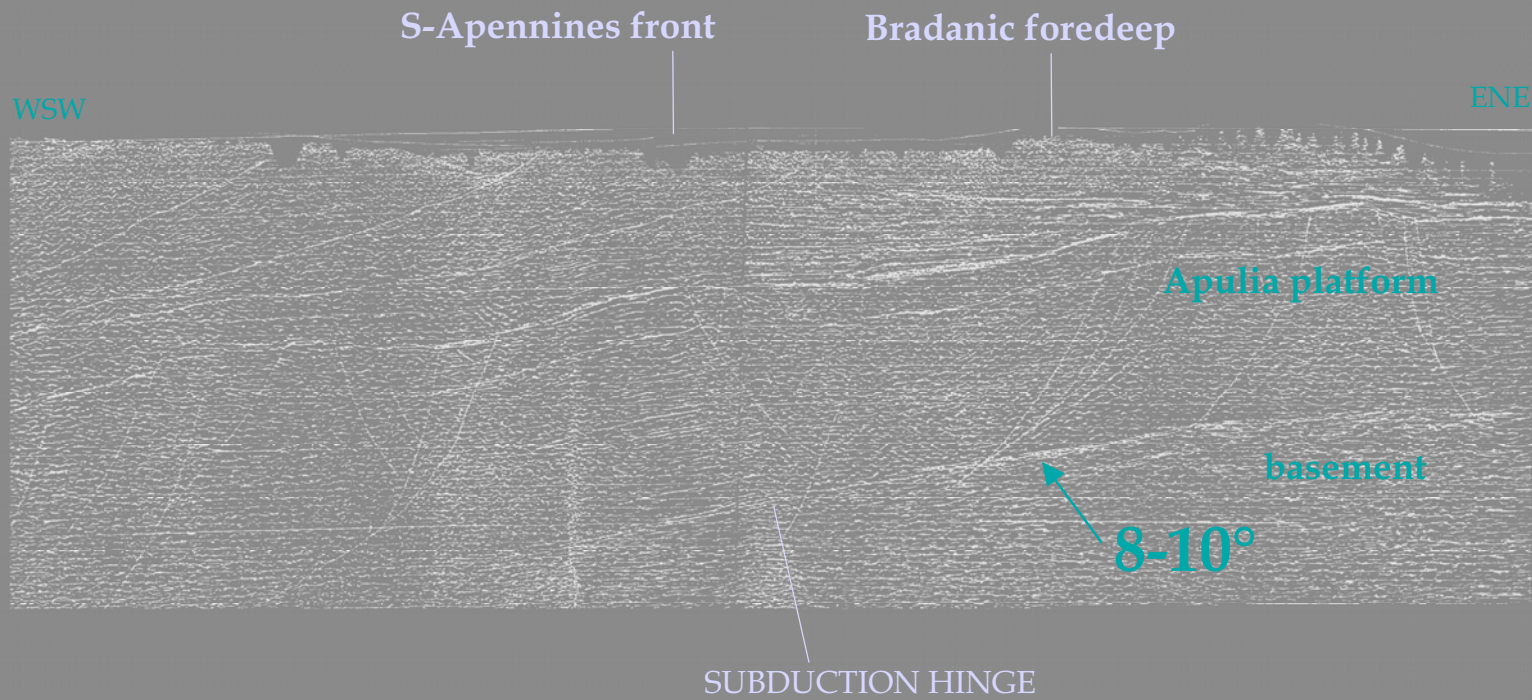




REGIONAL MONOCLINE



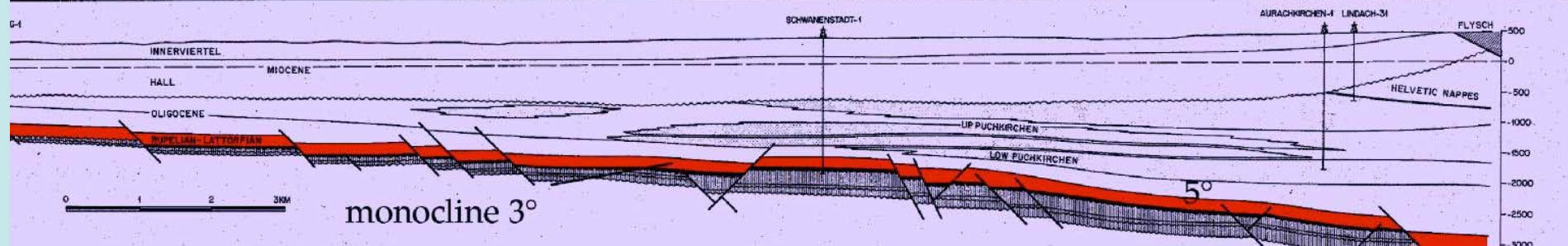
Southern Apennines



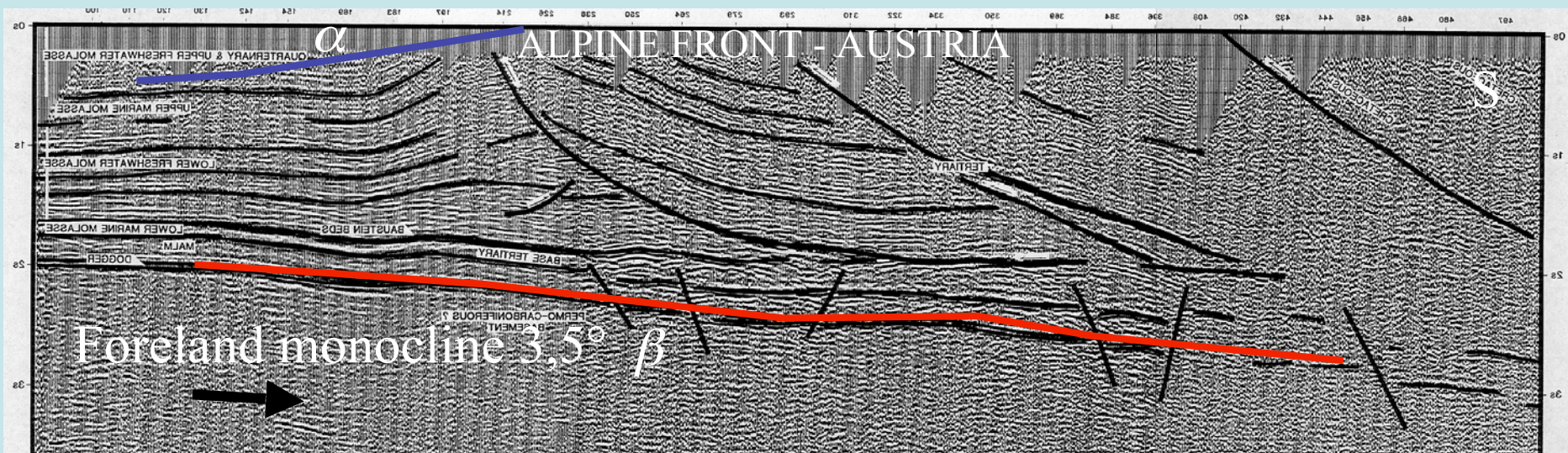
ALPS

AUSTRIA MOLASSE BASIN

G. Kittler and R. Neumayer
 Rohol-Aufsuchungs Ges. M.B.H.

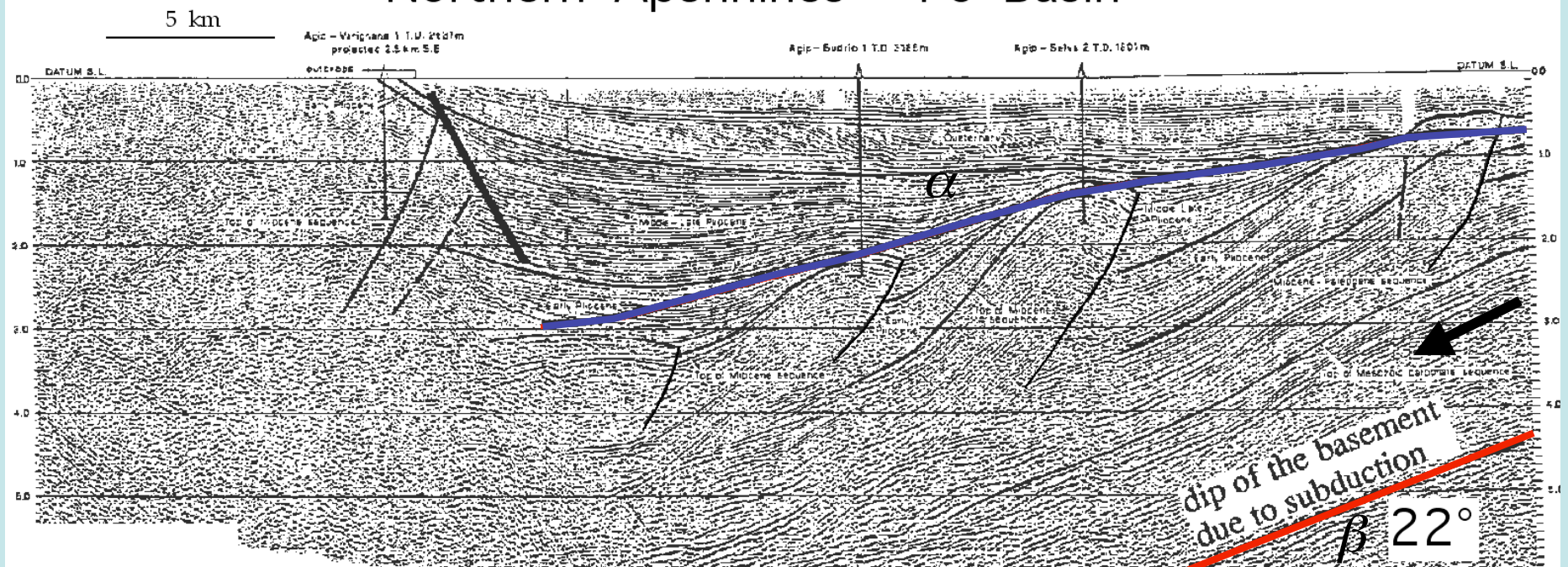


Kittler & Neumayer (1983)

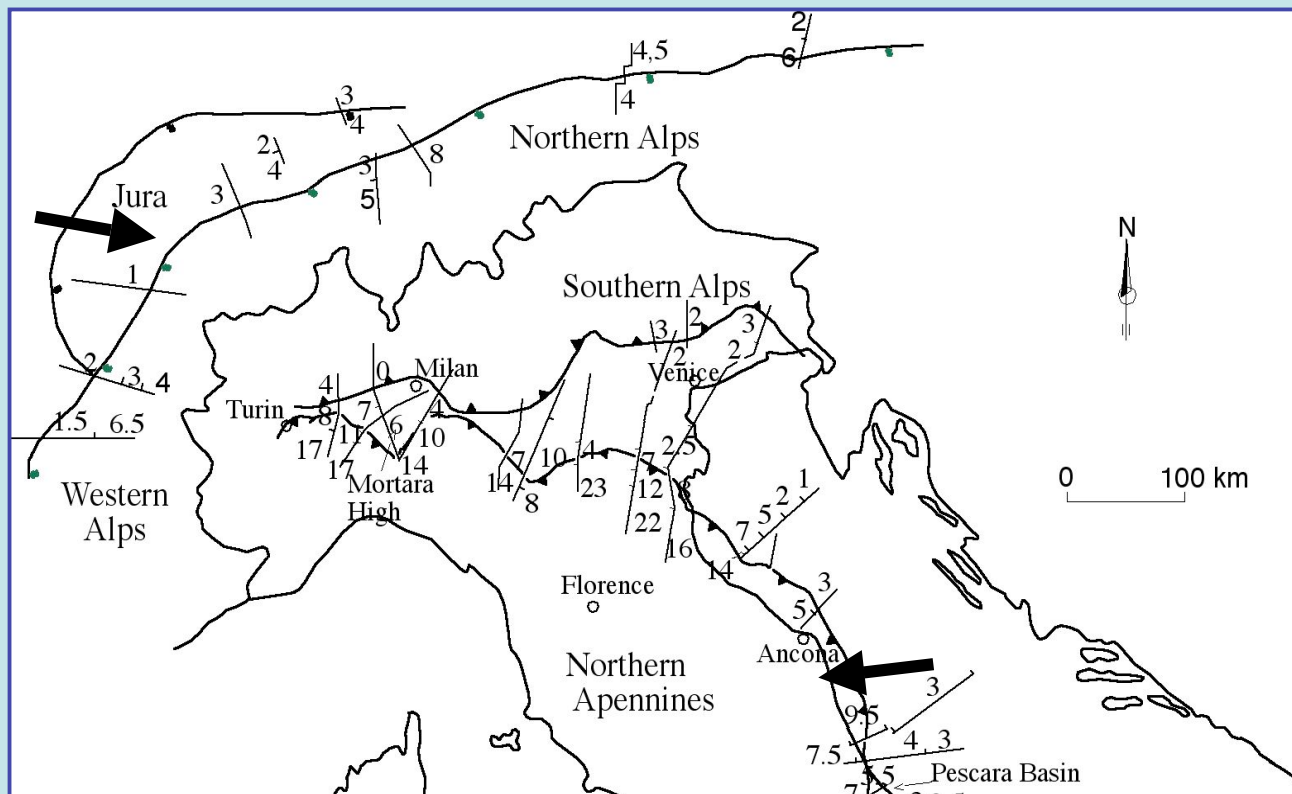


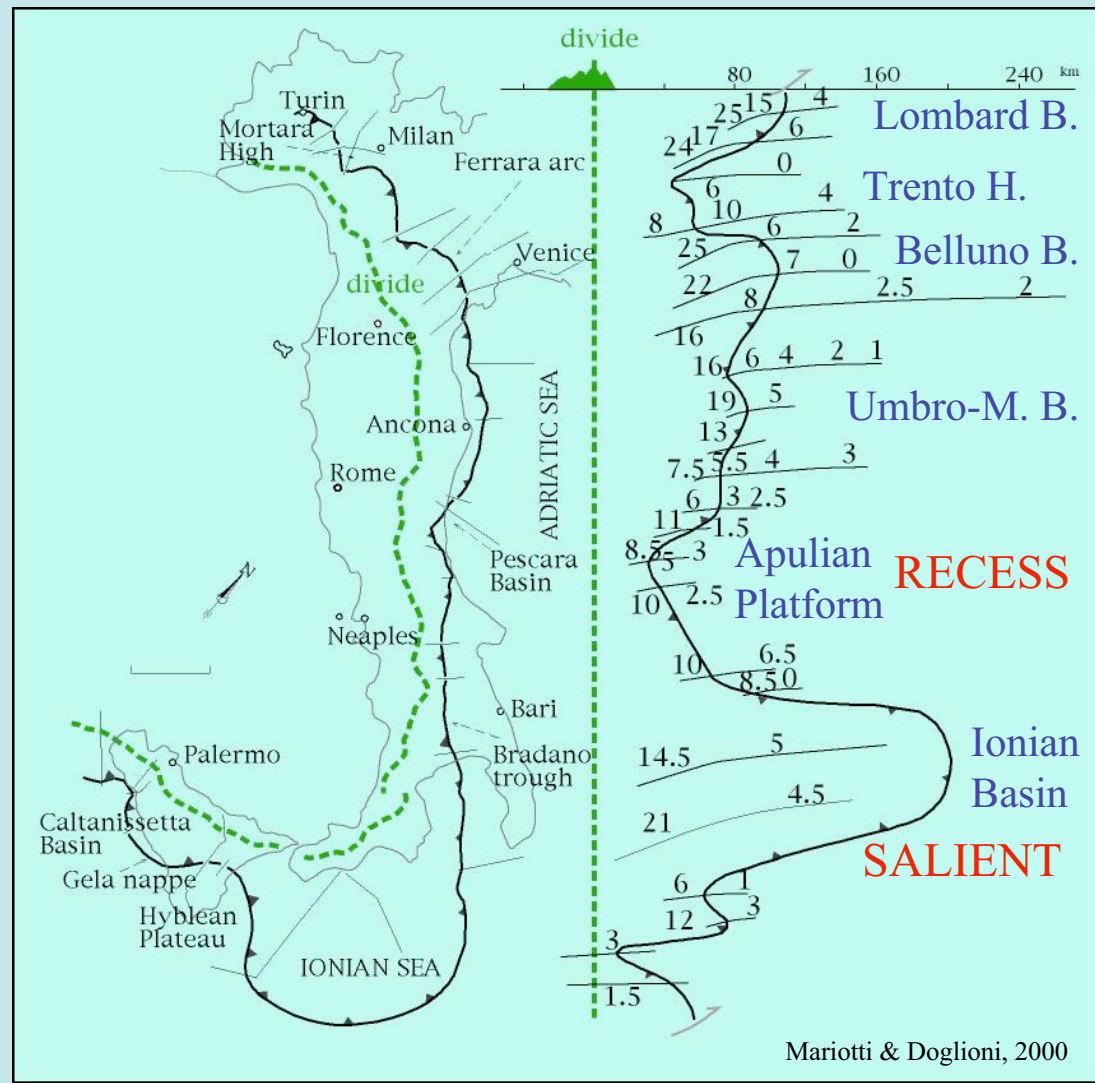
Bachmann & Koch, 1983

Northern Apennines - Po Basin

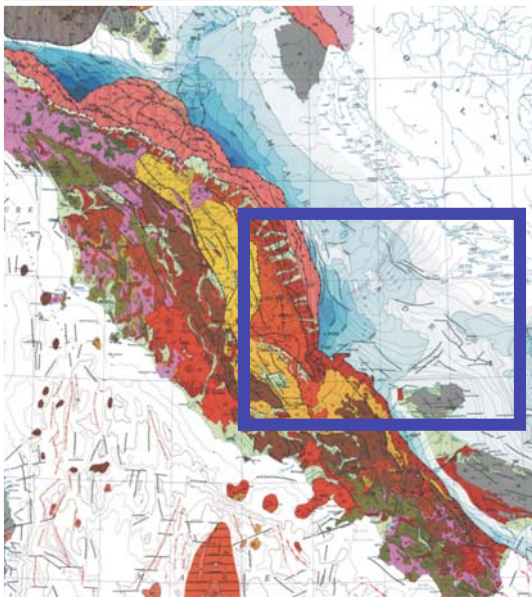
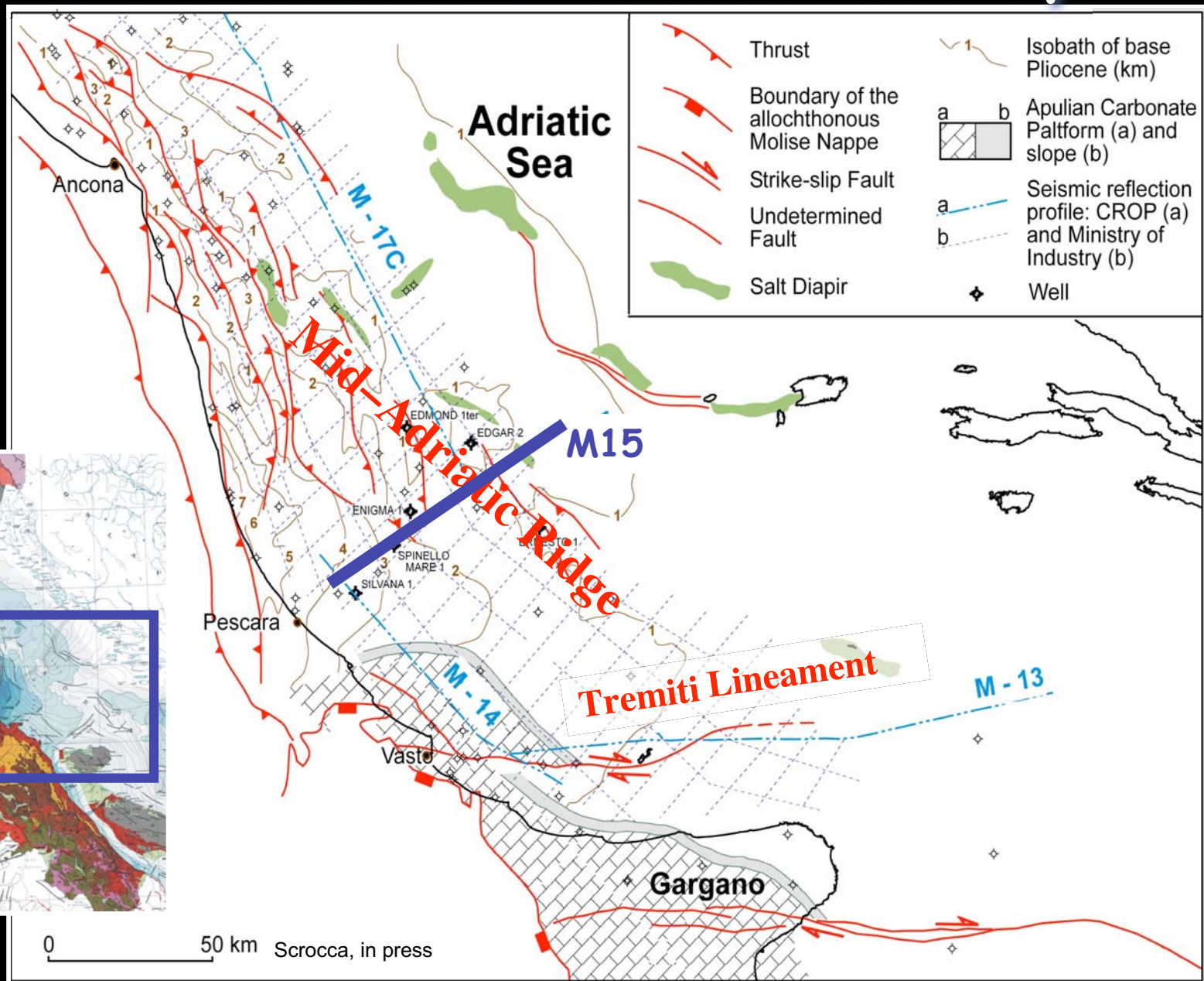


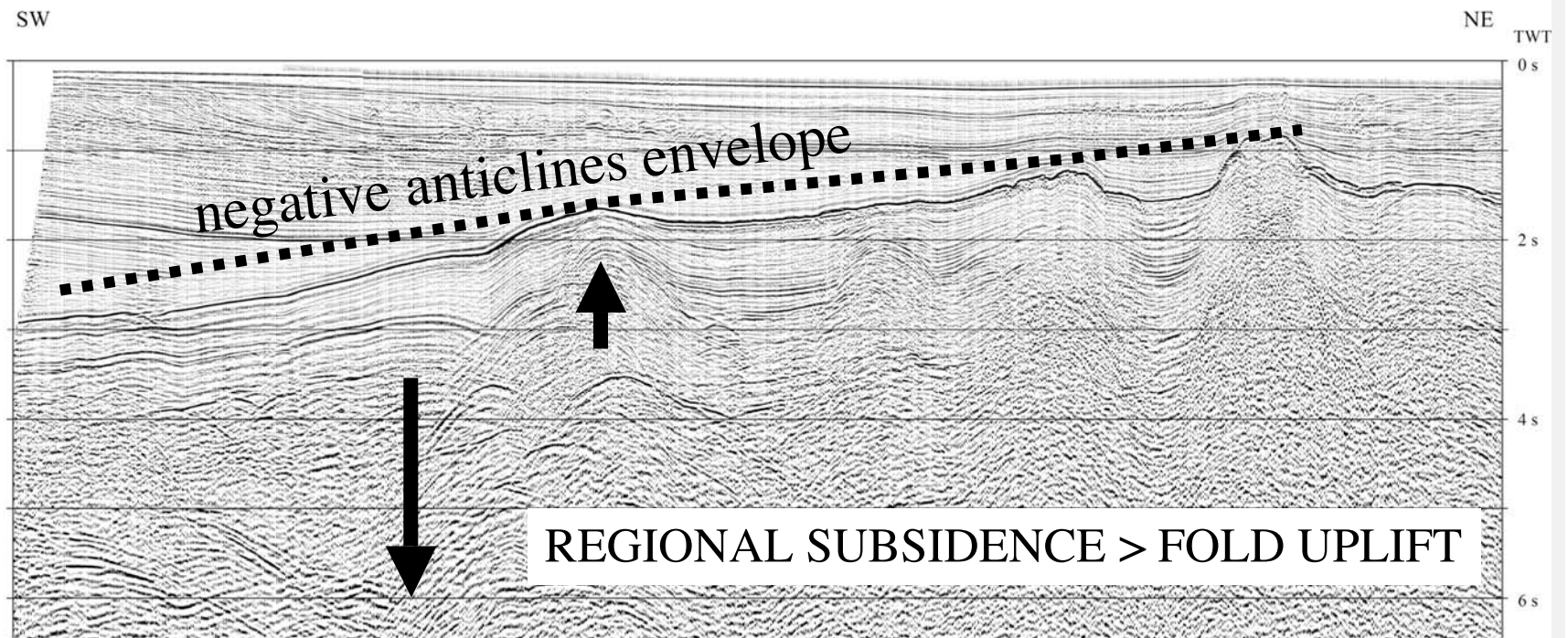
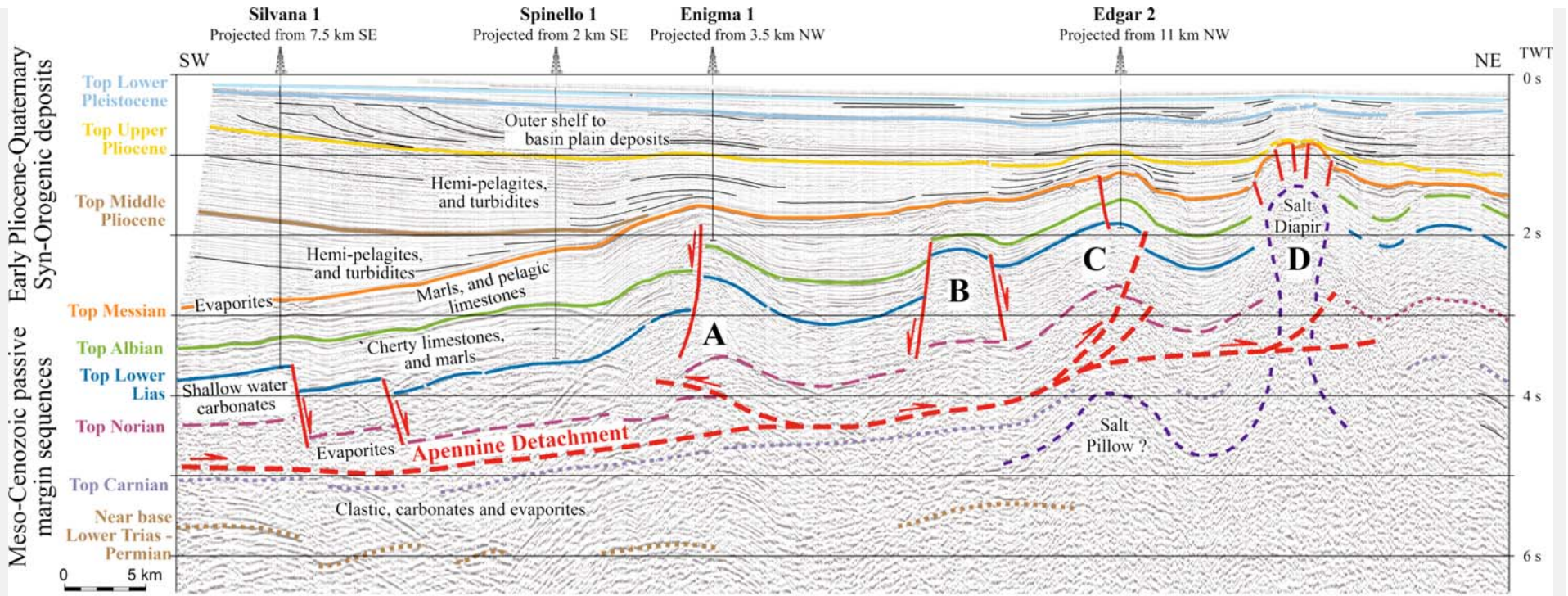
Pieri, 1983

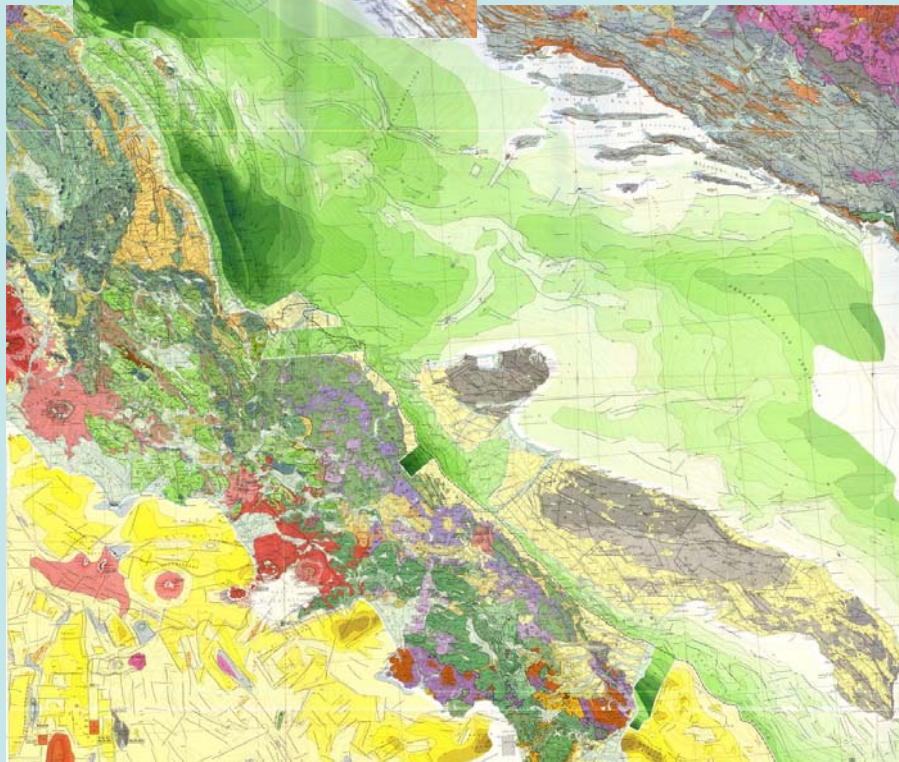




Central Adriatic Sea: Structural Map

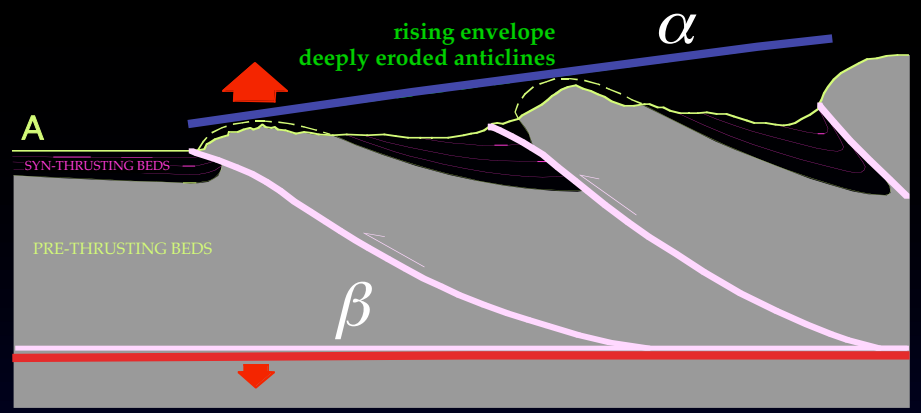




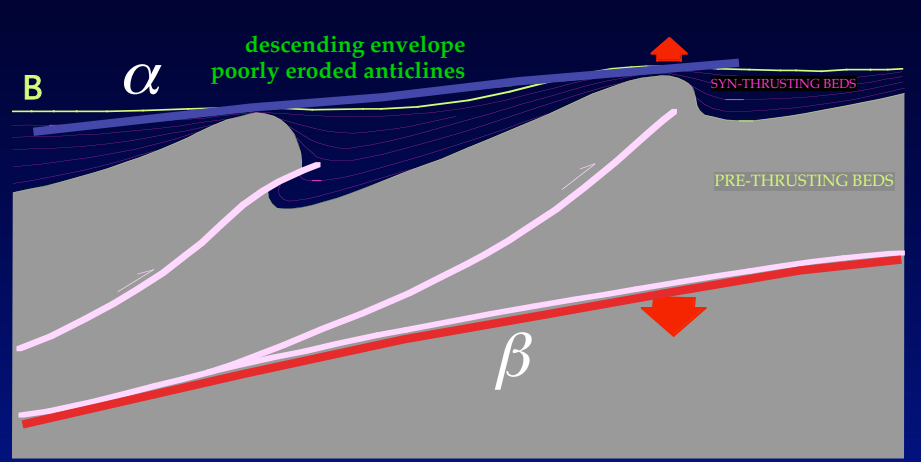


Alps - Dinarides

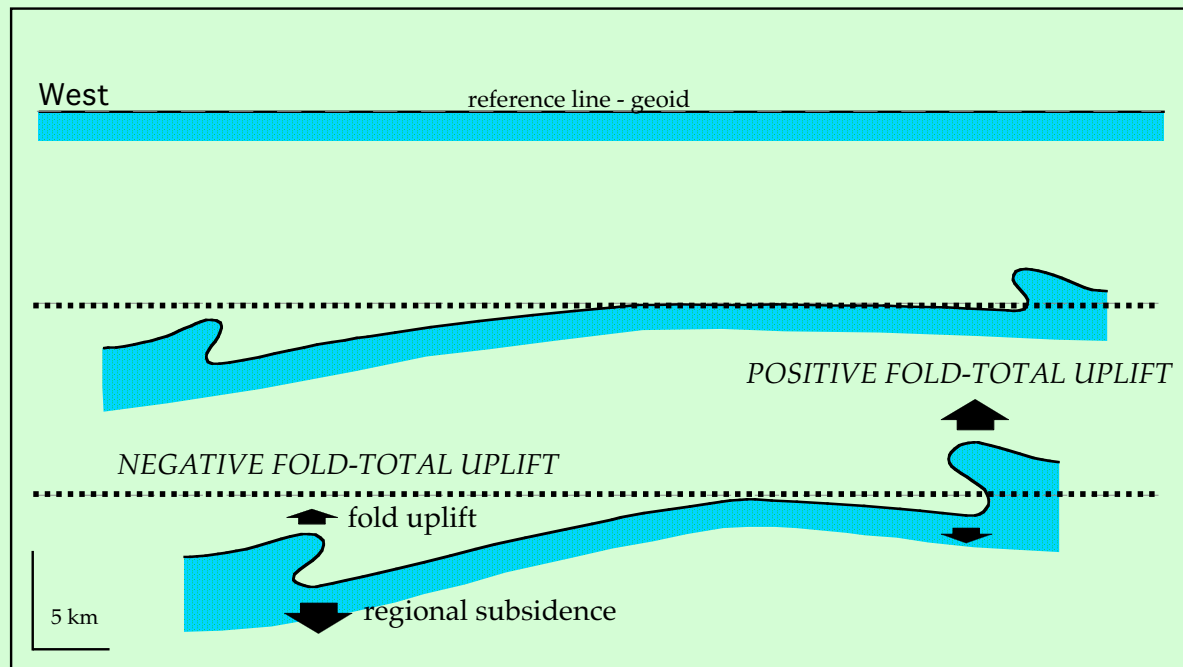
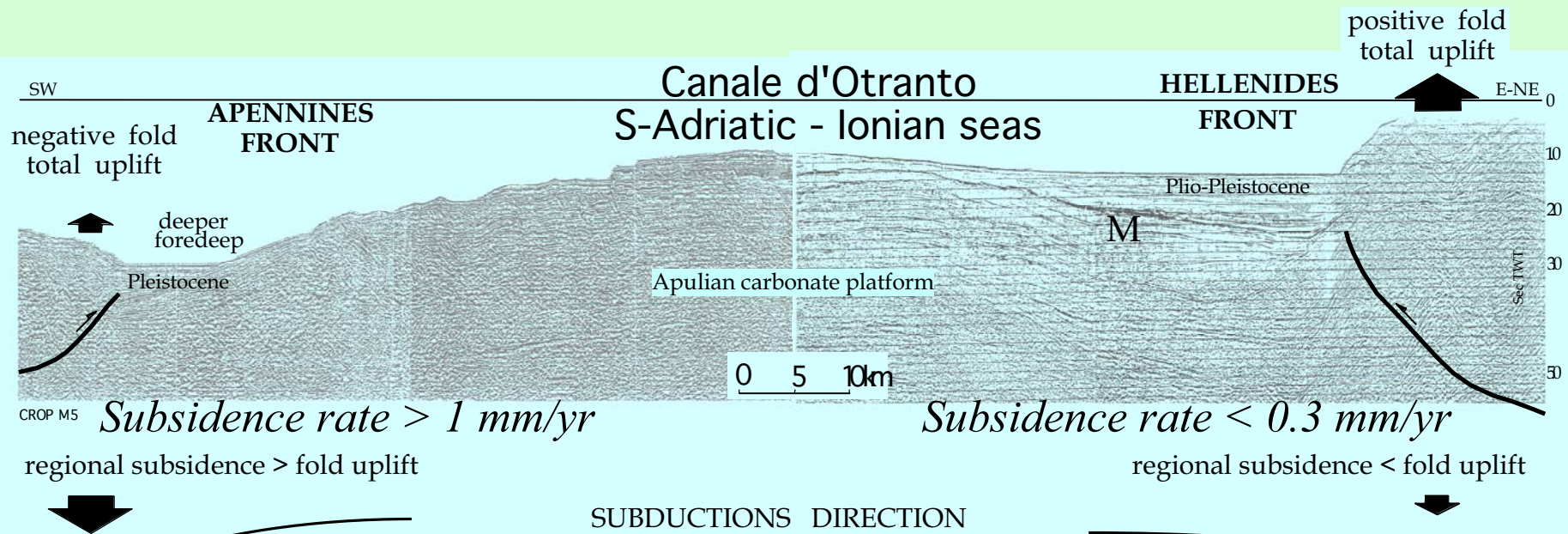
FOLD UPLIFT > REGIONAL SUBSIDENCE

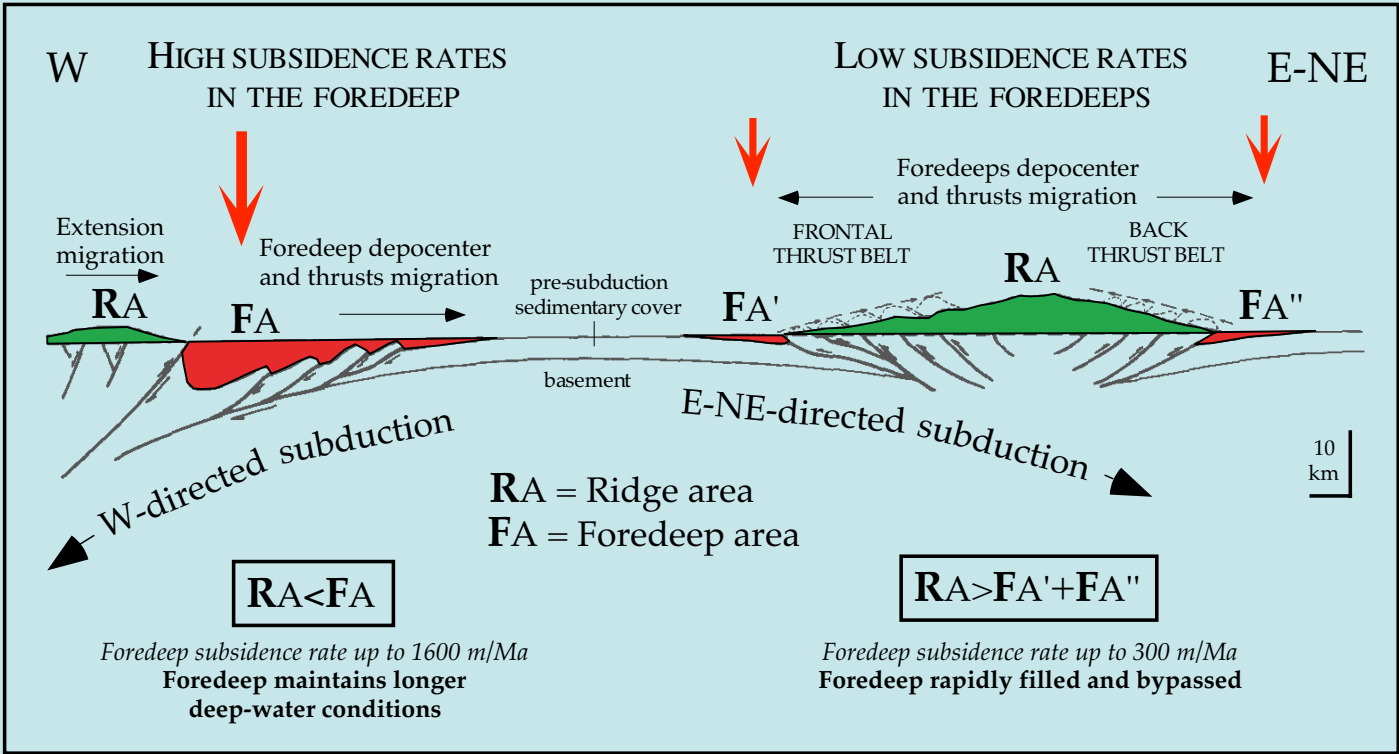


FOLD UPLIFT < REGIONAL SUBSIDENCE



Apennines



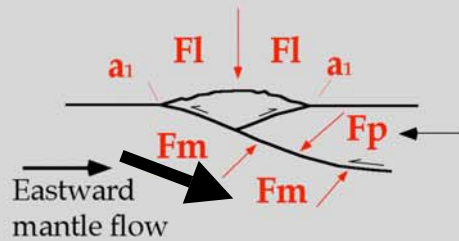


FOREDEEPS

EAST-DIRECTED SUBDUCTION

shallow foredeep

subsidence rates: 0-200 m/m.y.



Fm = mantle push
Fl = lithostatic load
Fp = upper plate push

a = foredeep basal dip

$a_2 > a_1$

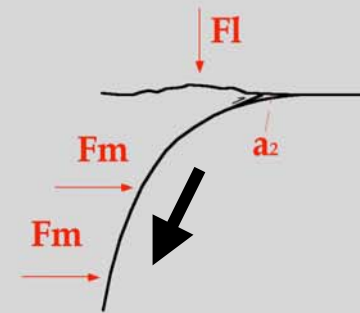
4 km Oligocene base

8 km Pliocene base

WEST-DIRECTED SUBDUCTION

deep foredeep

subsidence rates: 800-1000 m/m.y.

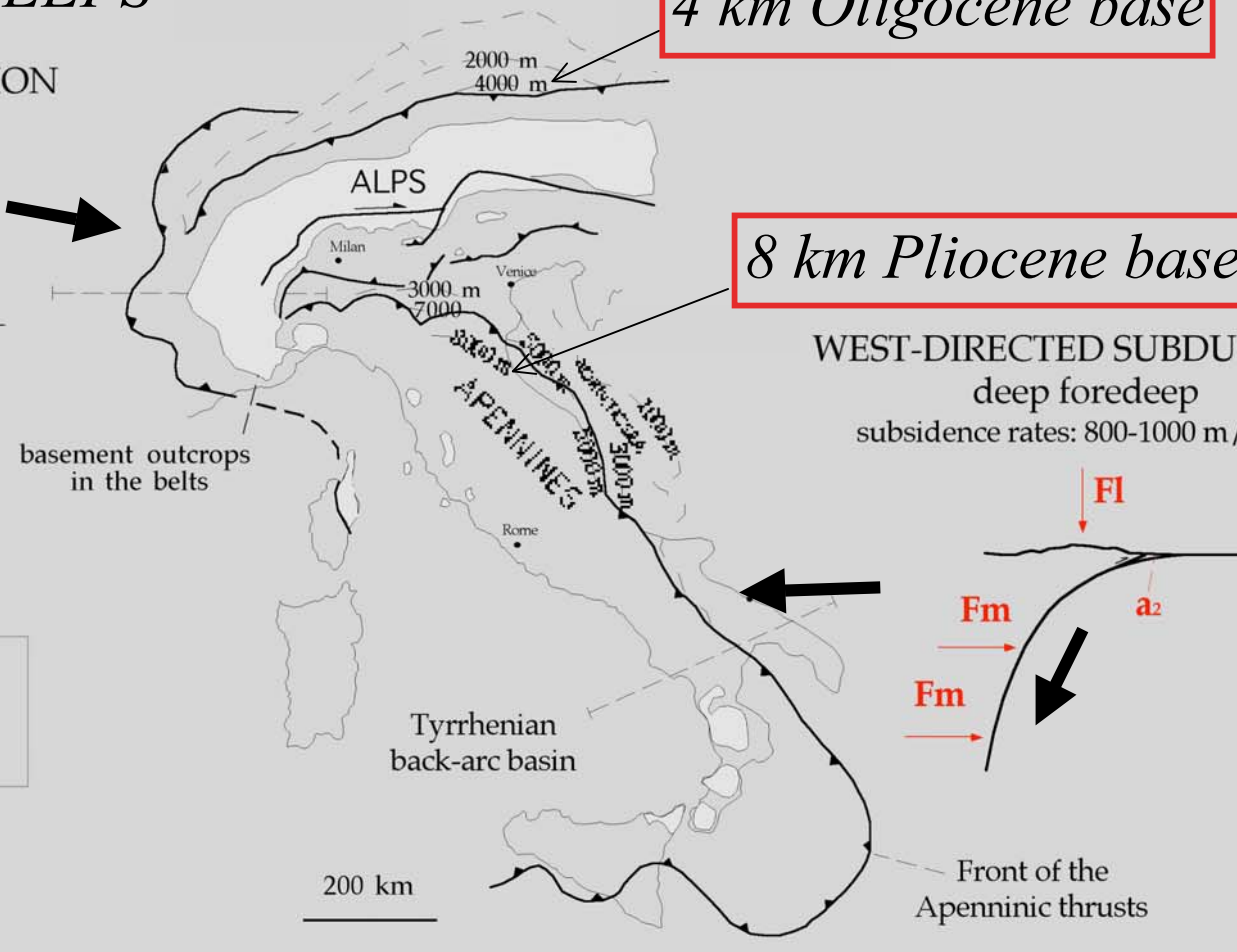


basement outcrops
in the belts

Tyrrhenian
back-arc basin

Front of the
Apenninic thrusts

200 km

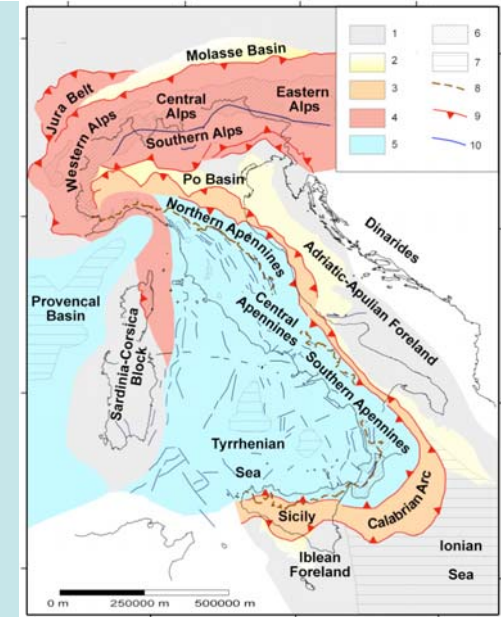


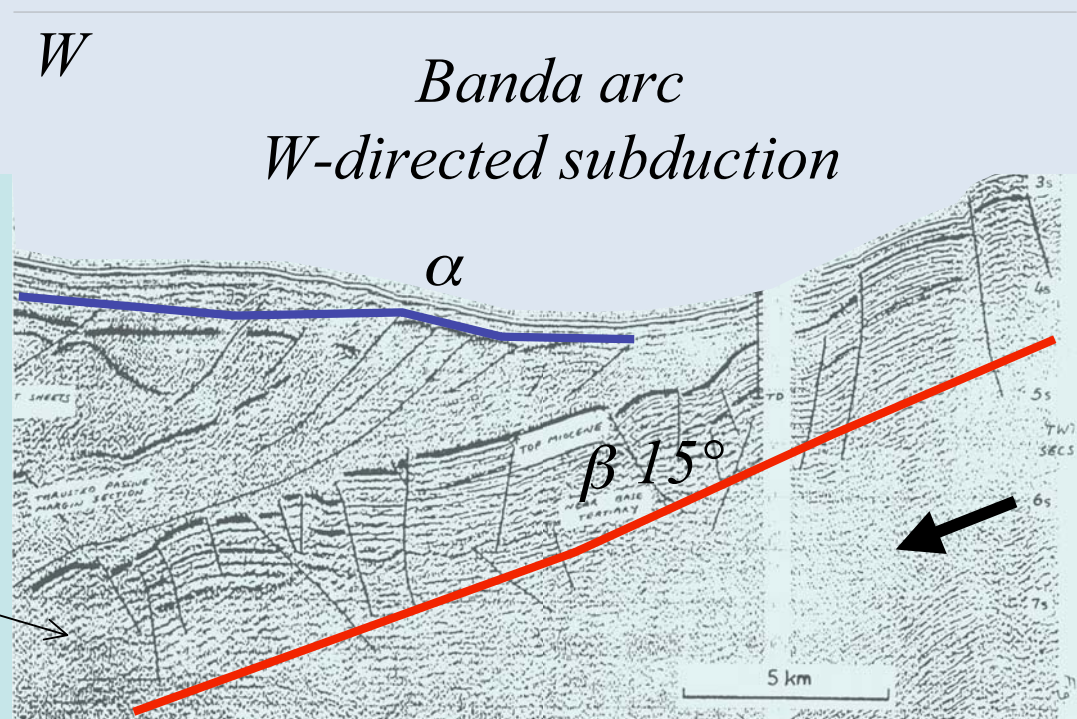
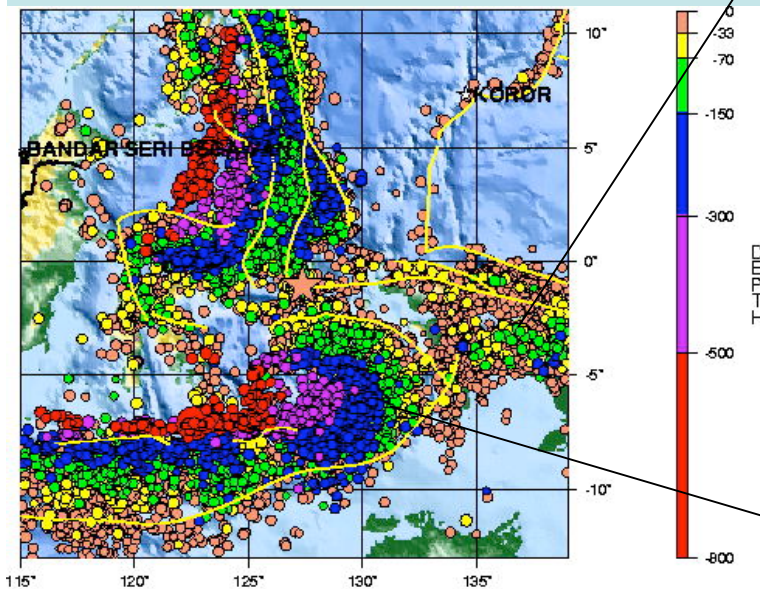
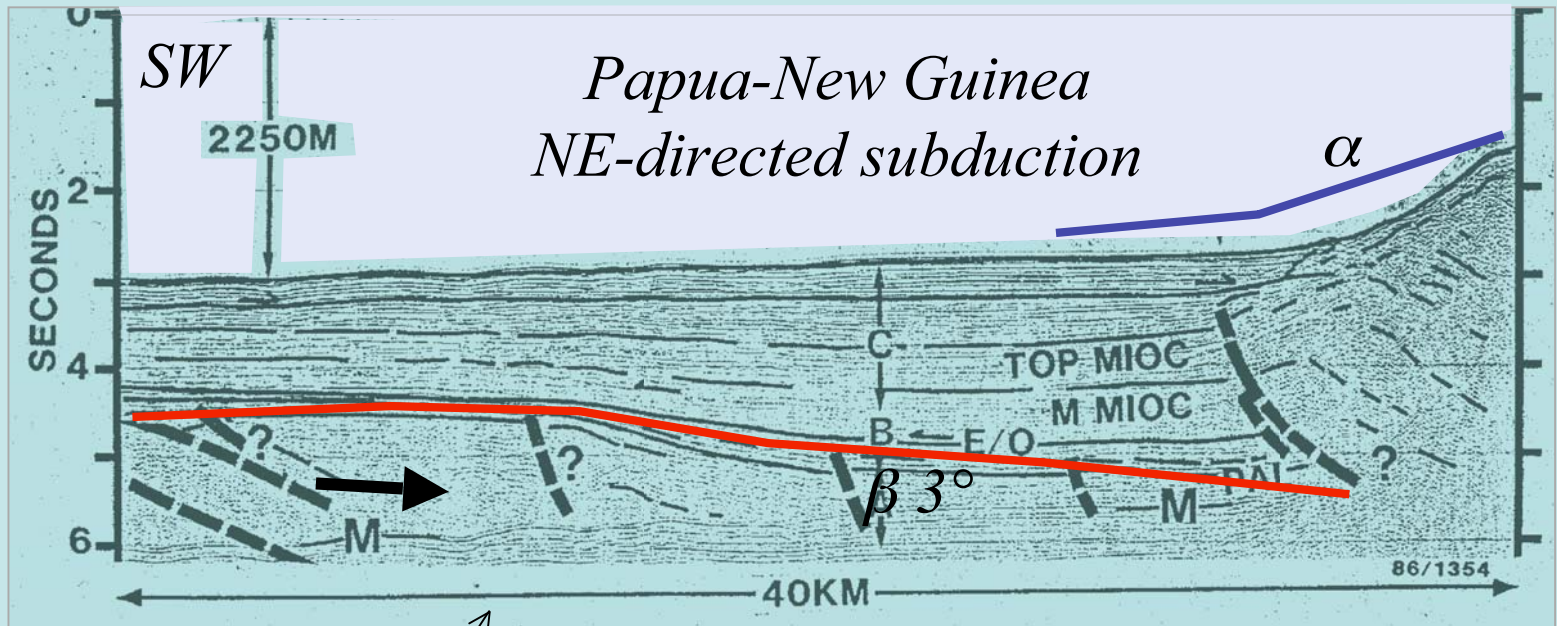
ALPS

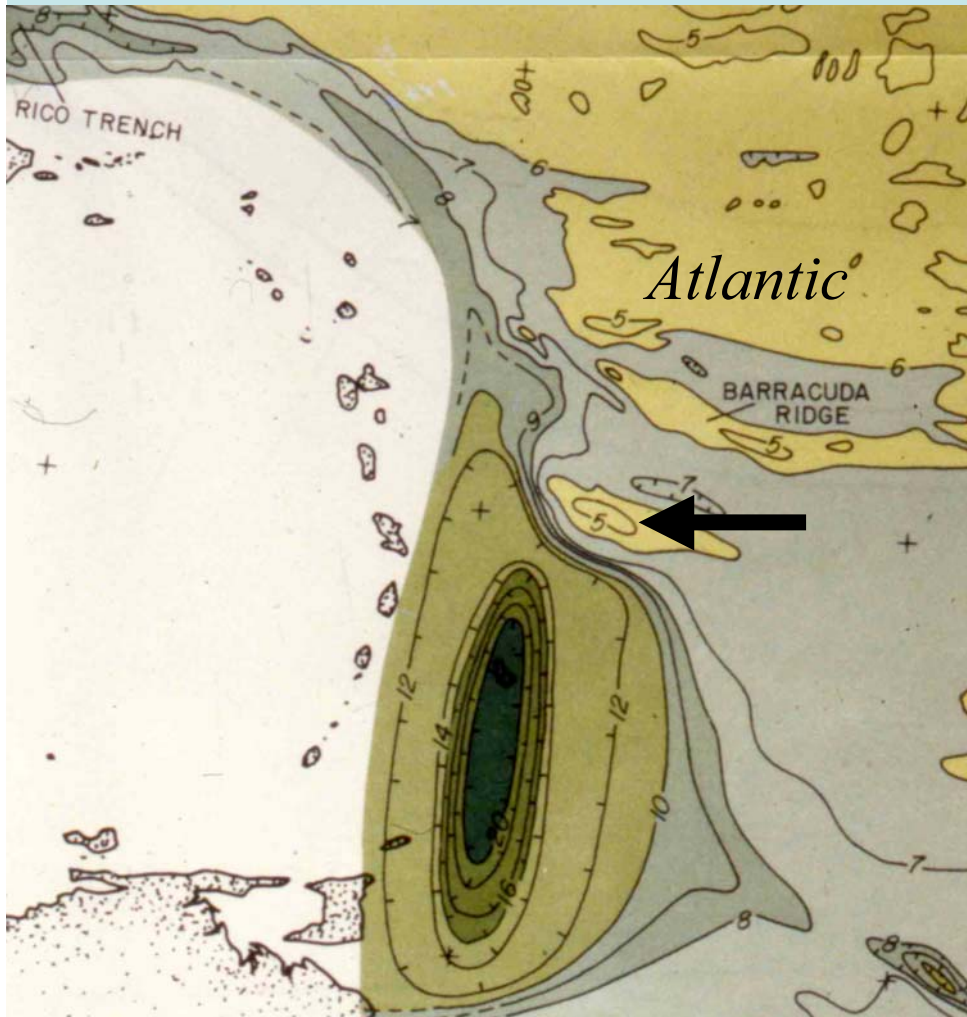
- double vergence
- high topography
- shallow foredeep
- deep rocks involved
- shallow foreland monocline dip
- thickened crust under the belt
- thickened lithosphere
- no backarc basin

APENNINES

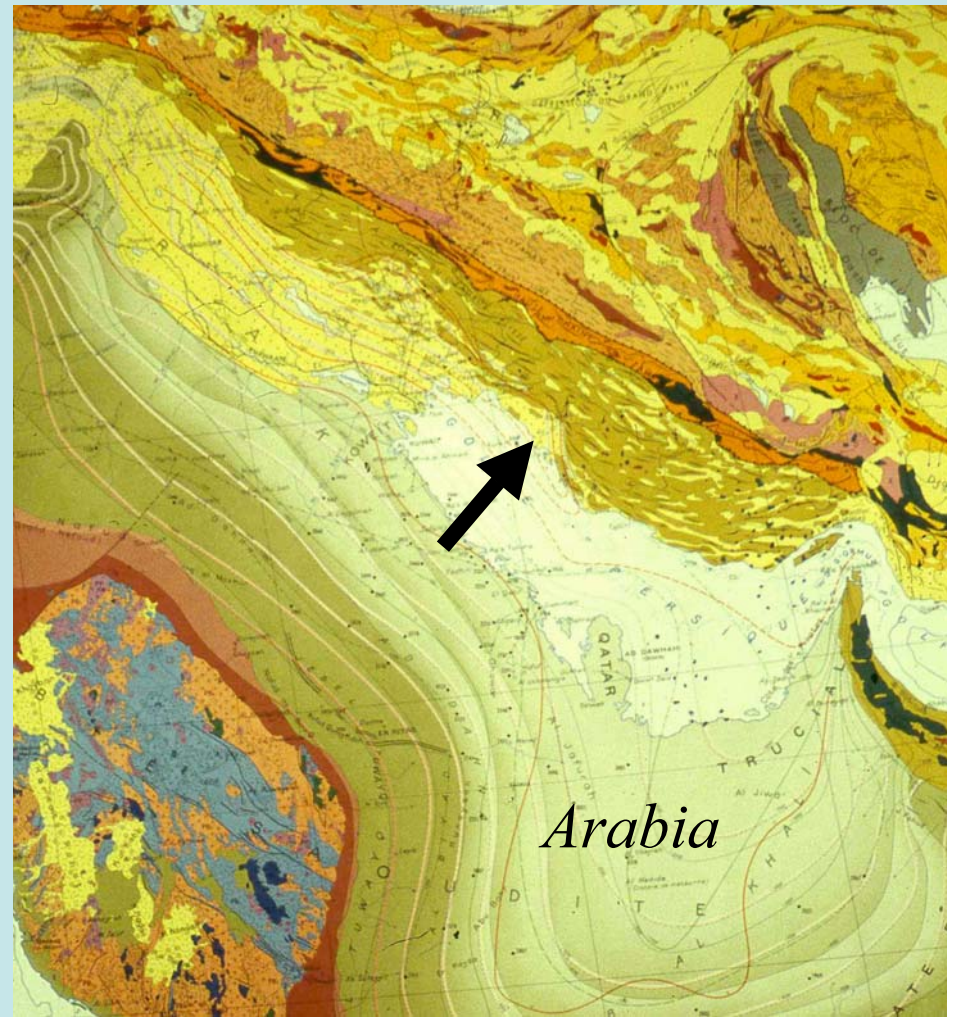
- single vergence
- low topography
- deep foredeep
- shallow rocks involved
- steep foreland monocline dip
- thinned crust under the belt
- shallow hangingwall asthenosphere
- widespread hangingwall extension & well developed backarc basin





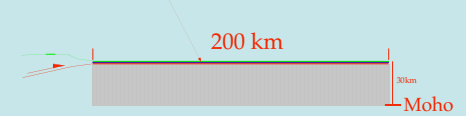


Barbados W-subduction

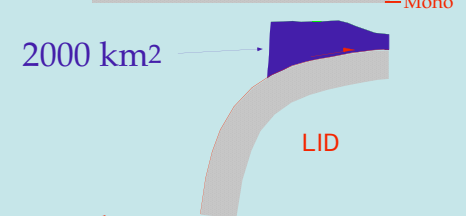
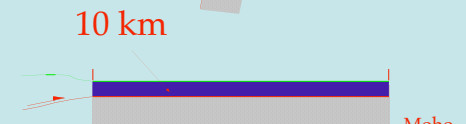
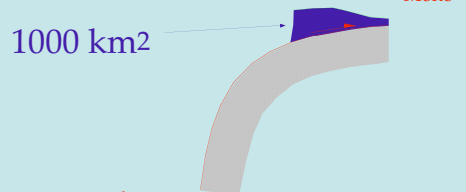
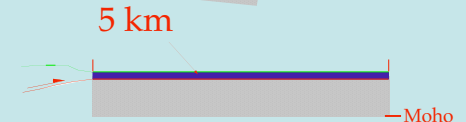
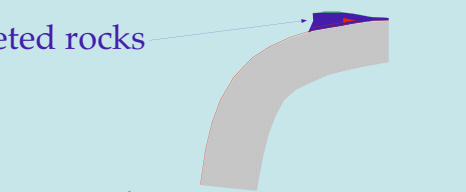


Zagros NE-subduction

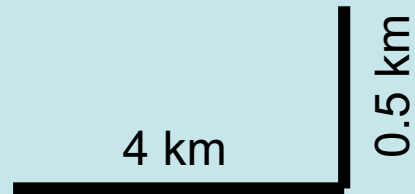
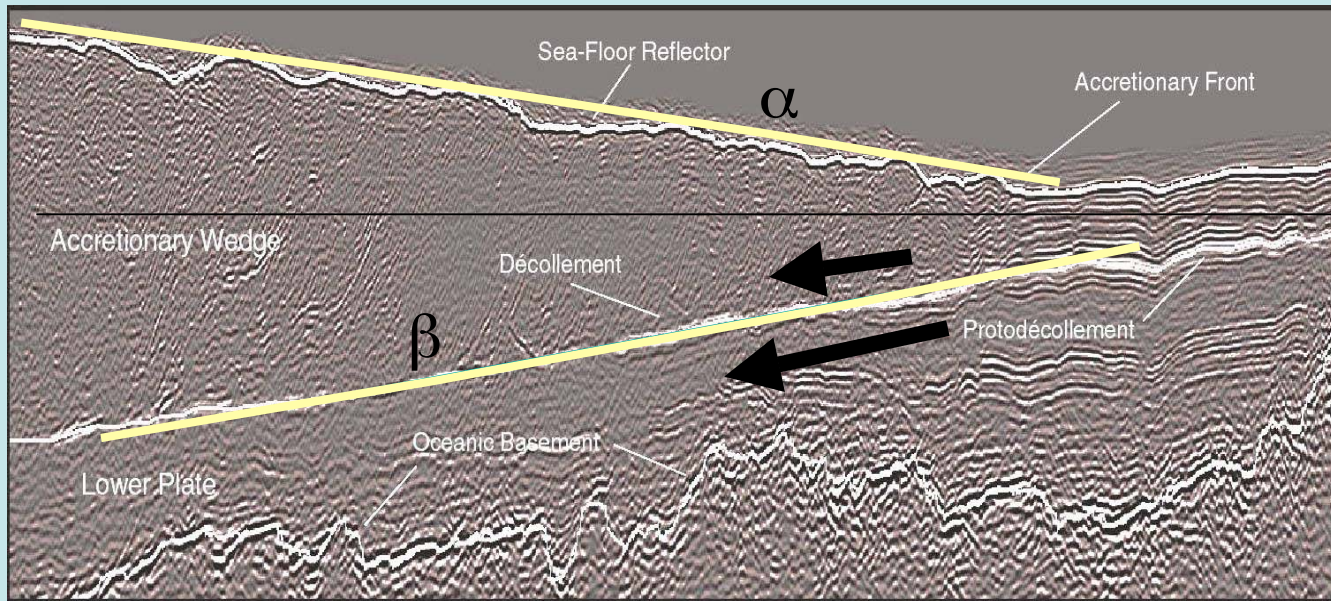
DECOLLEMENT DEPTH 2.5 km

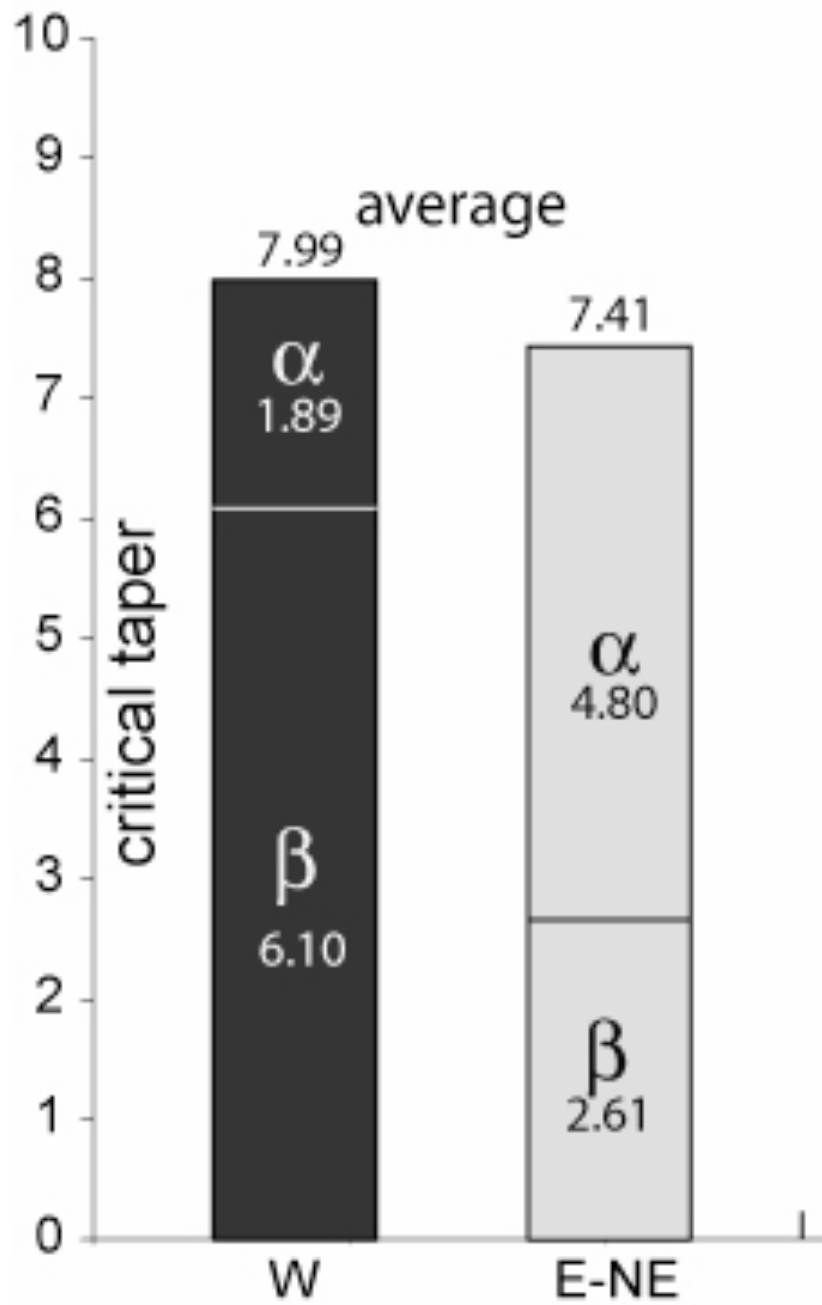


500 km² of accreted rocks
~10% erosion

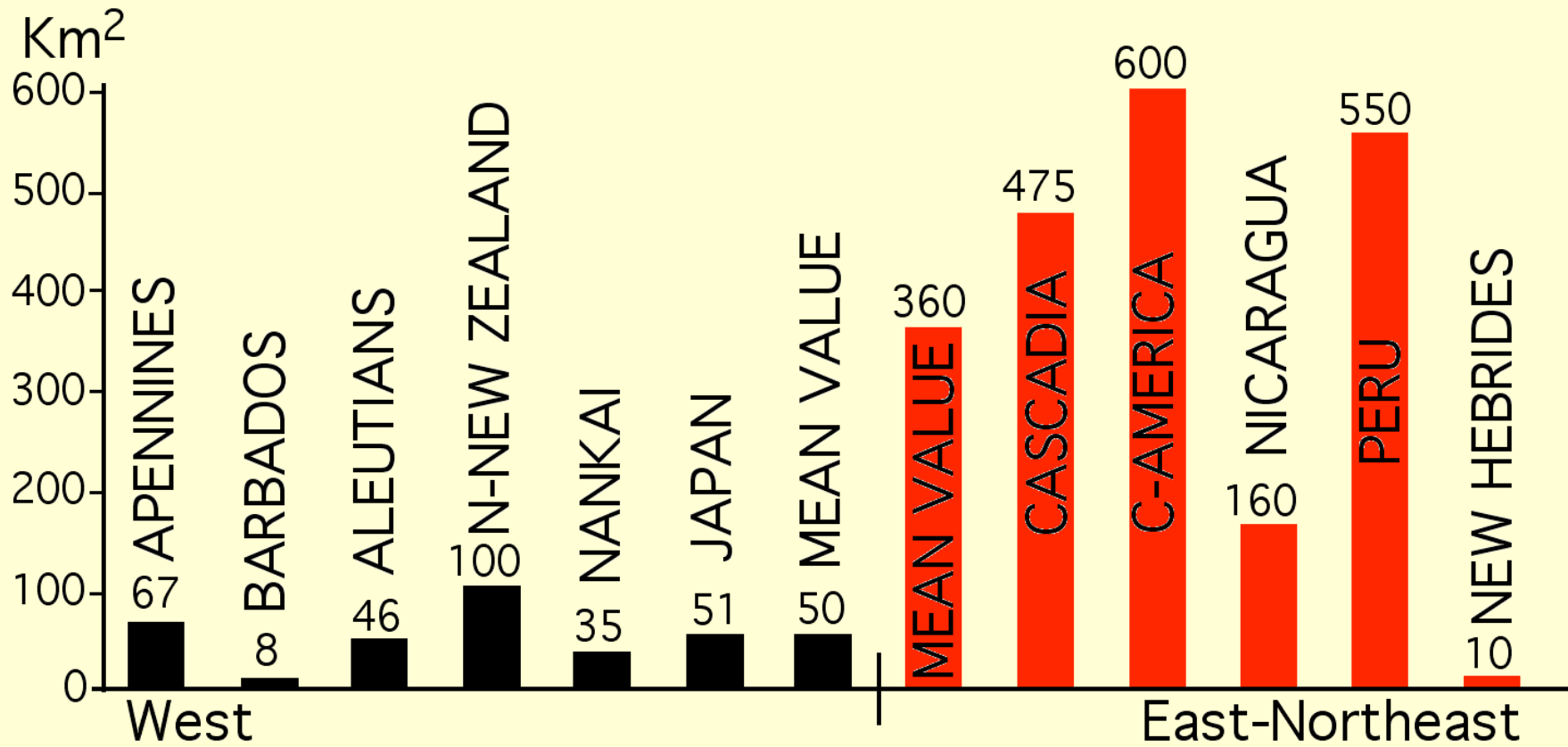


4000 km²
25km structural elevation
~50% erosion



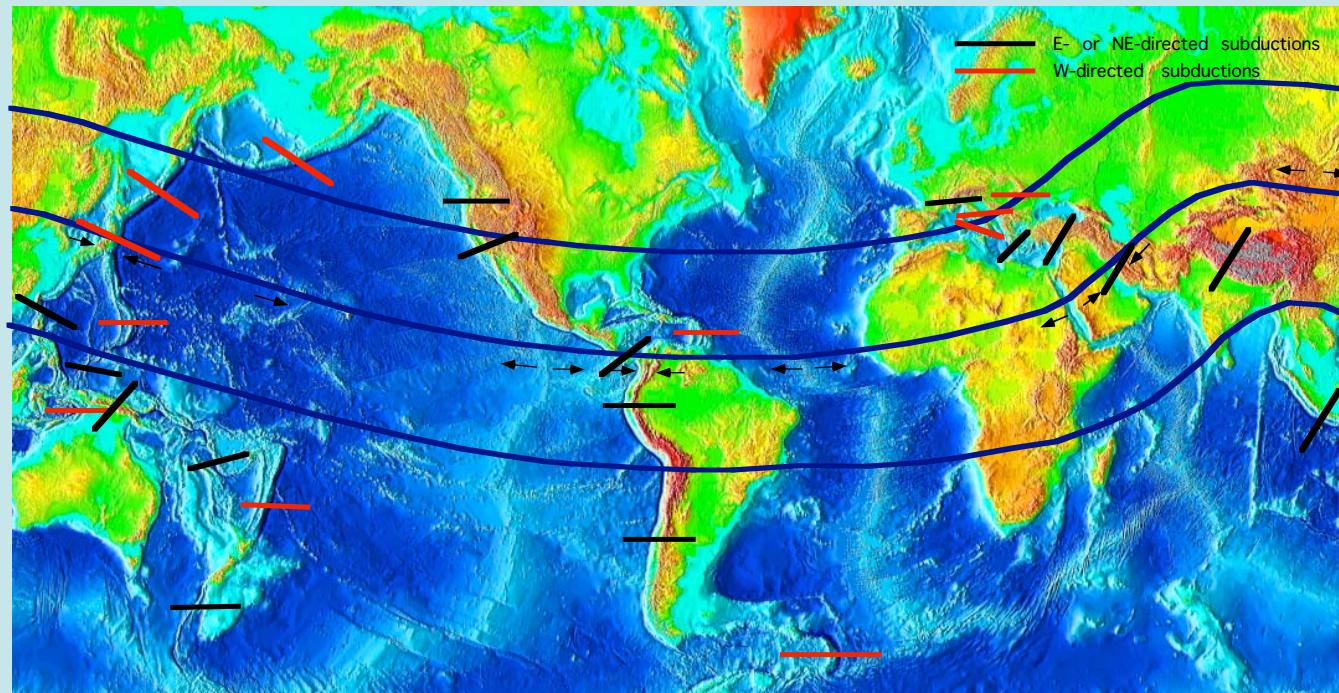
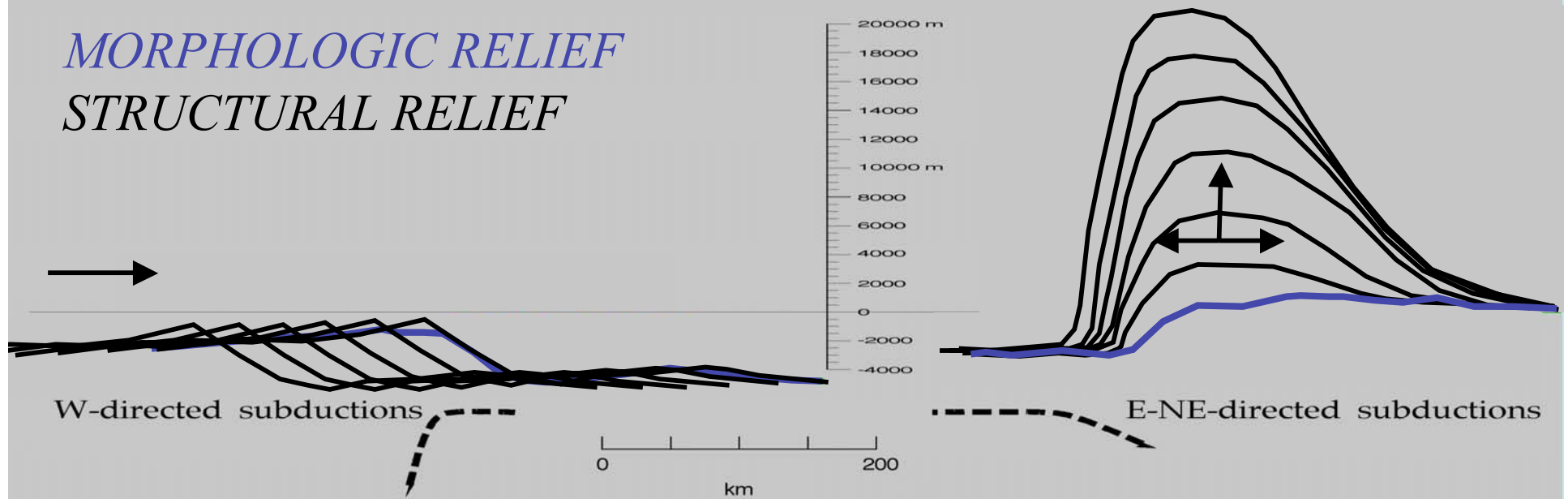


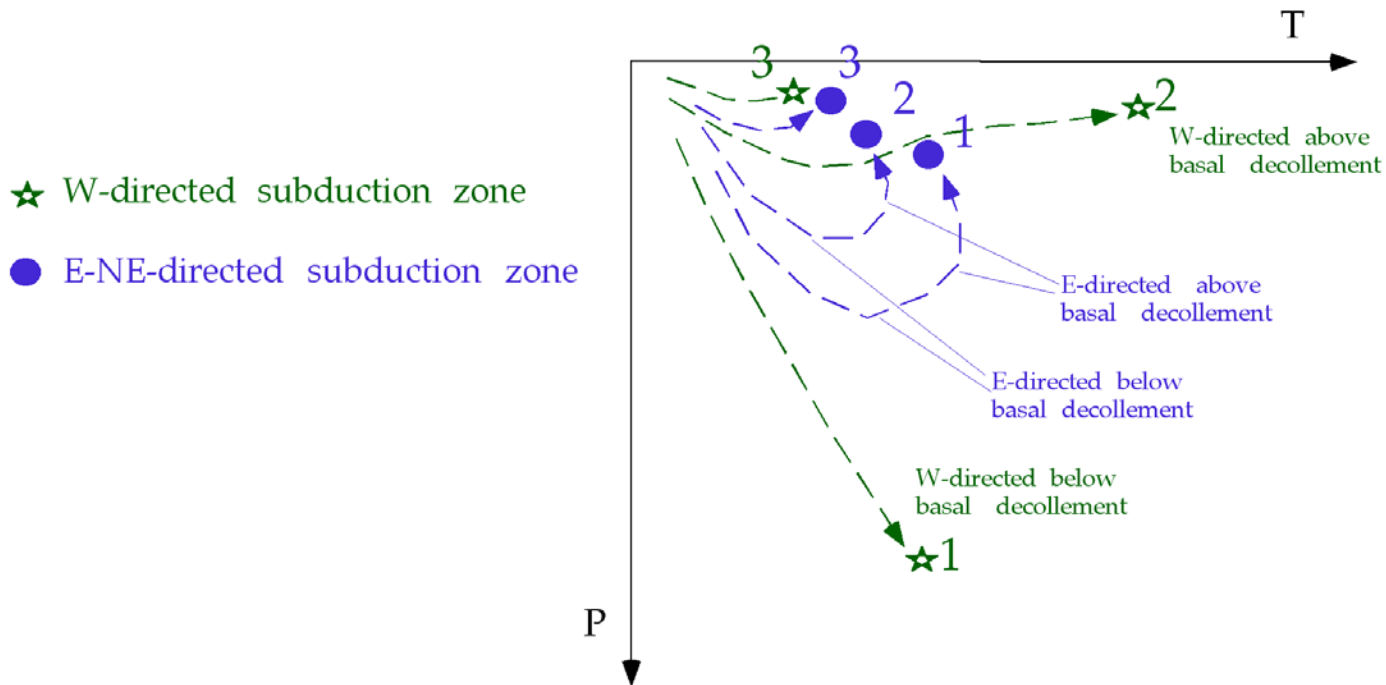
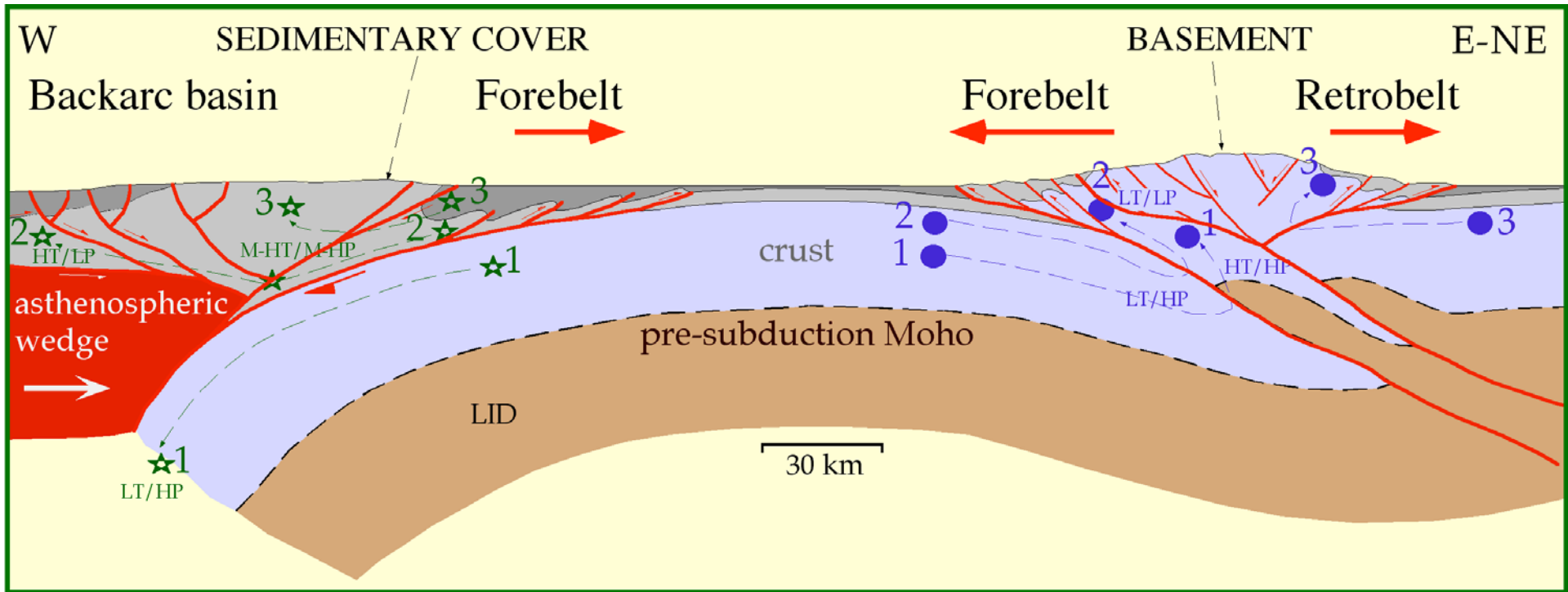
Lenci & Doglioni, 2006

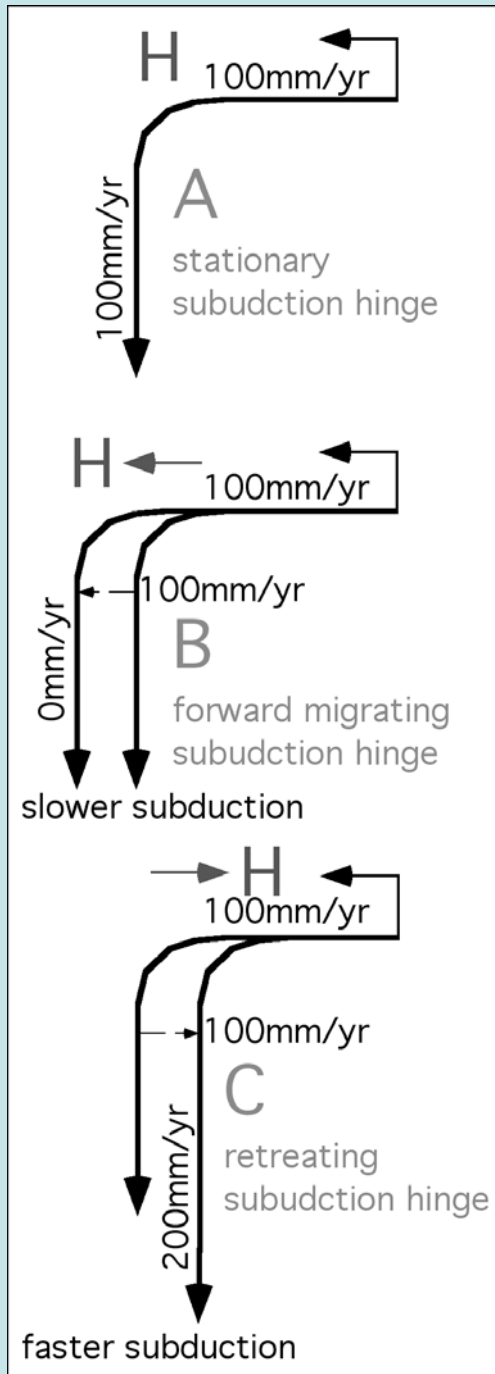


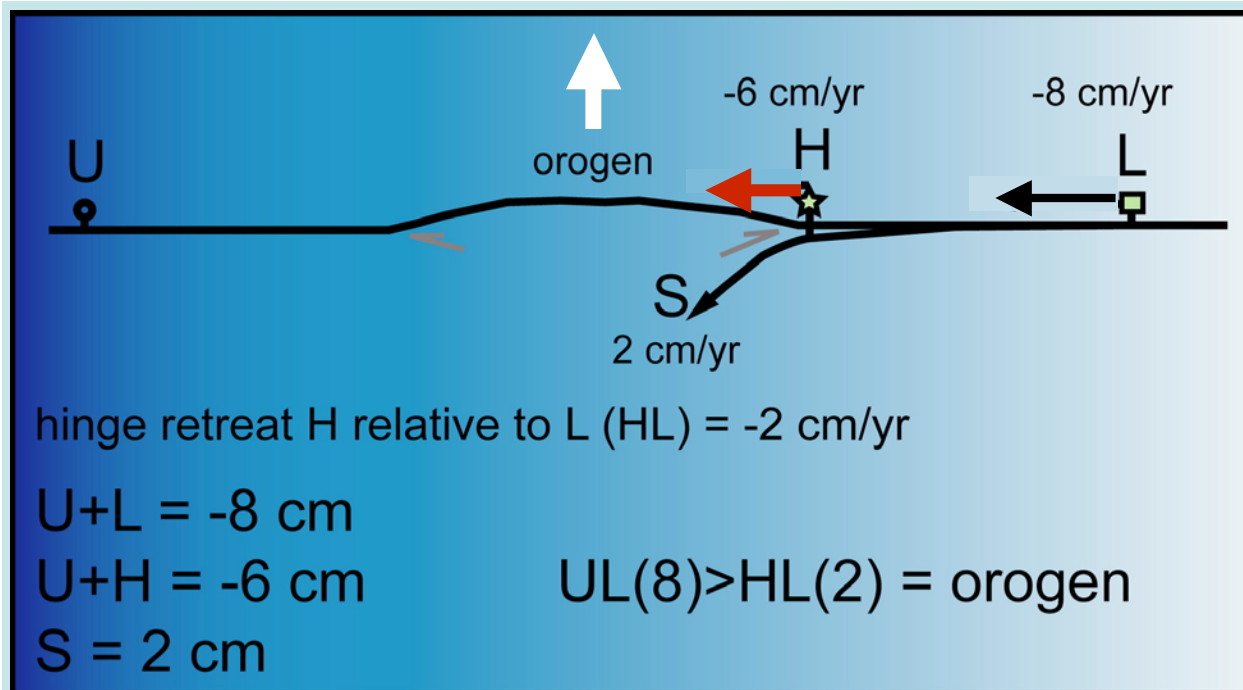
orogens area above sea level

MORPHOLOGIC RELIEF
STRUCTURAL RELIEF





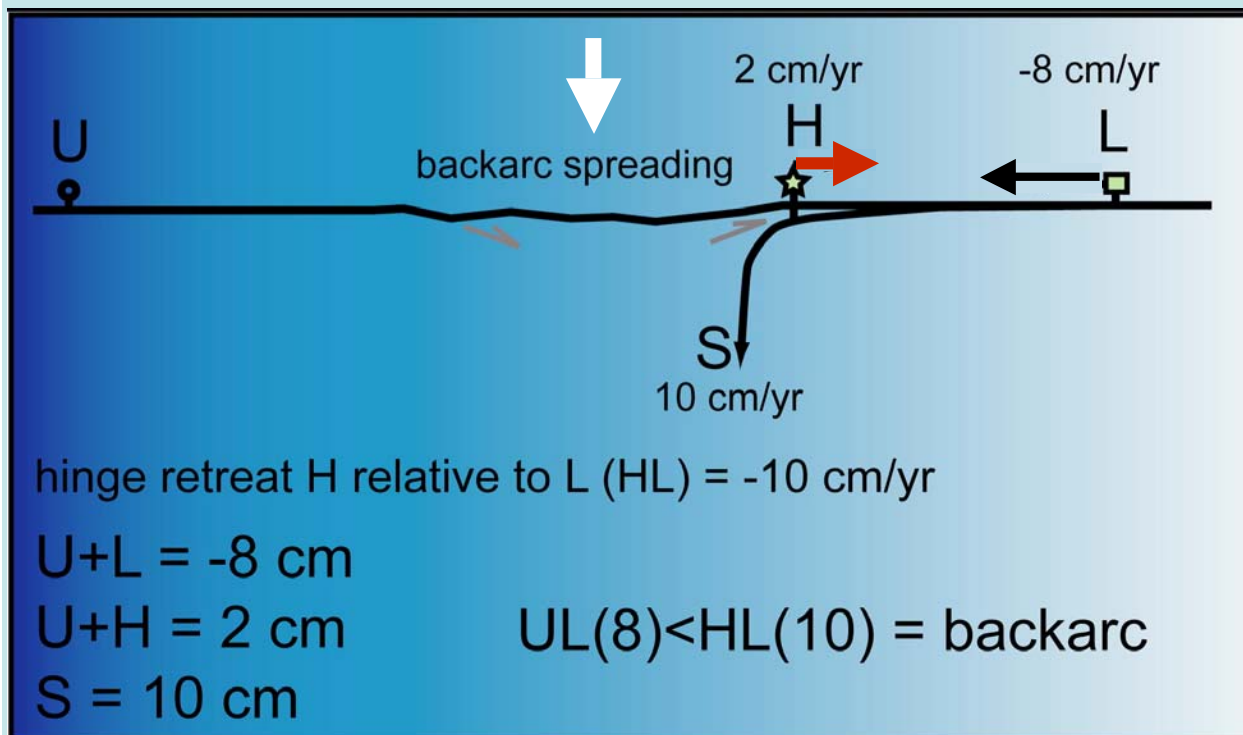




U=upper plate (fixed)
 H=subduction hinge
 L=lower plate
 S=subduction rate

Subduction < Convergence

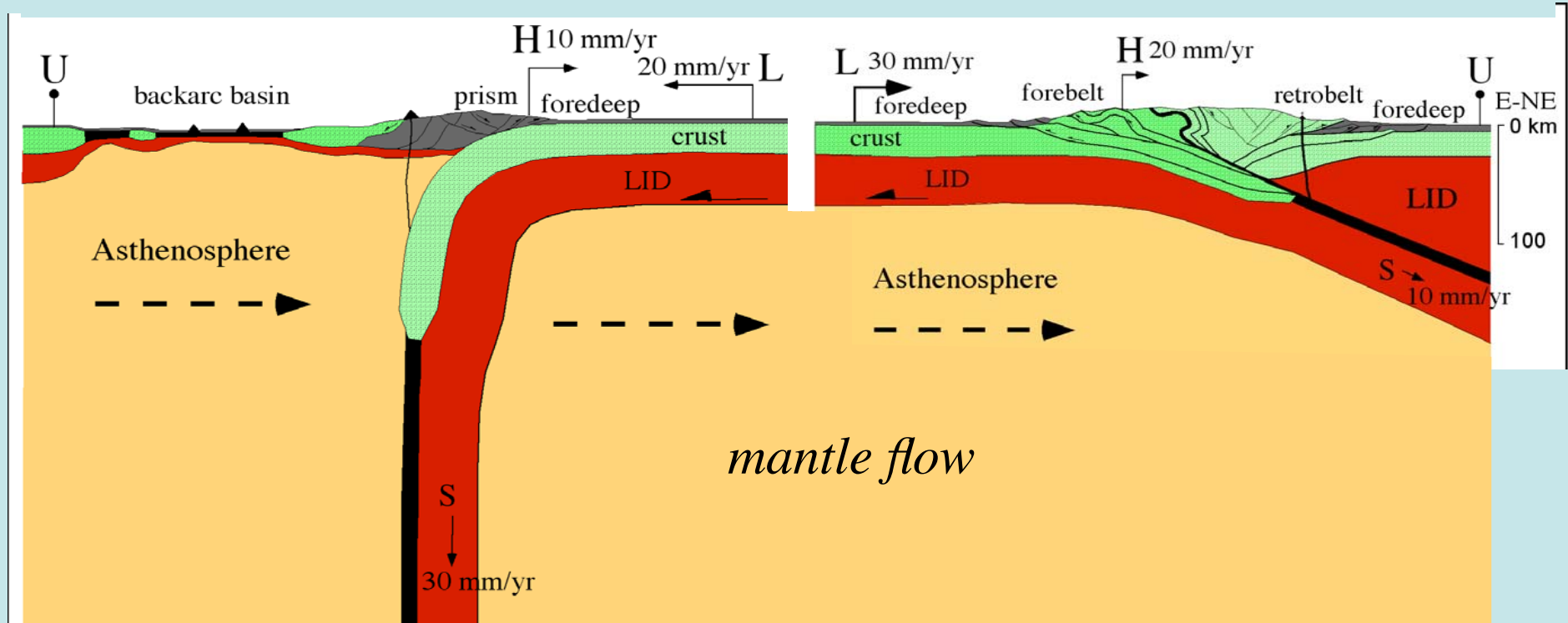
$$S = H - L$$

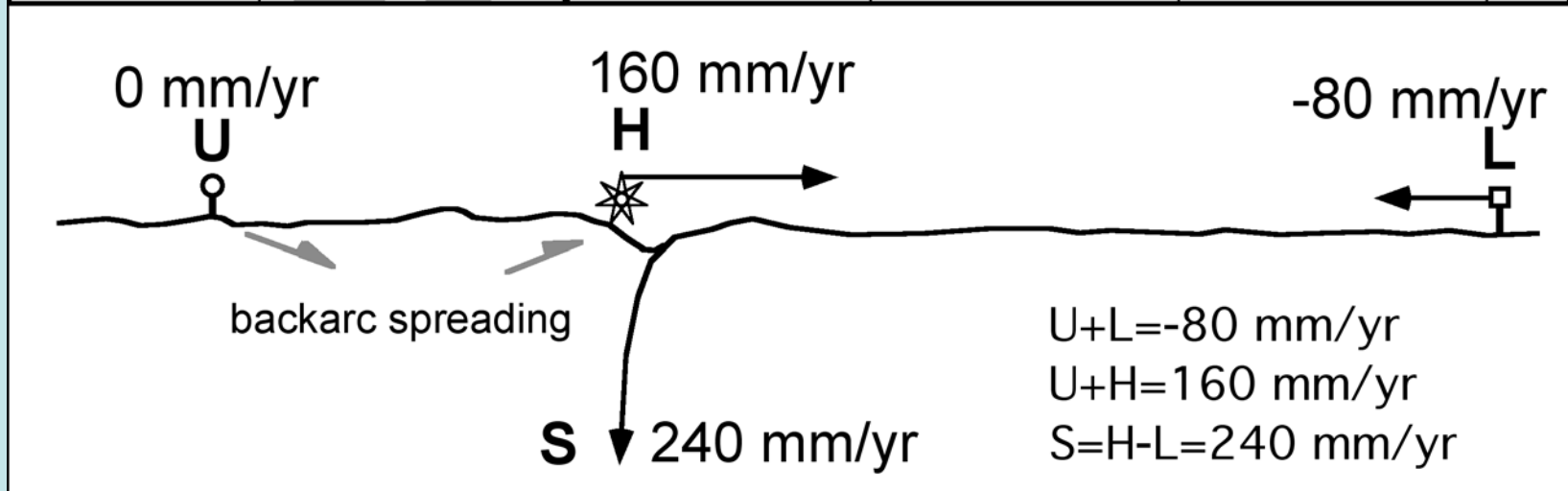
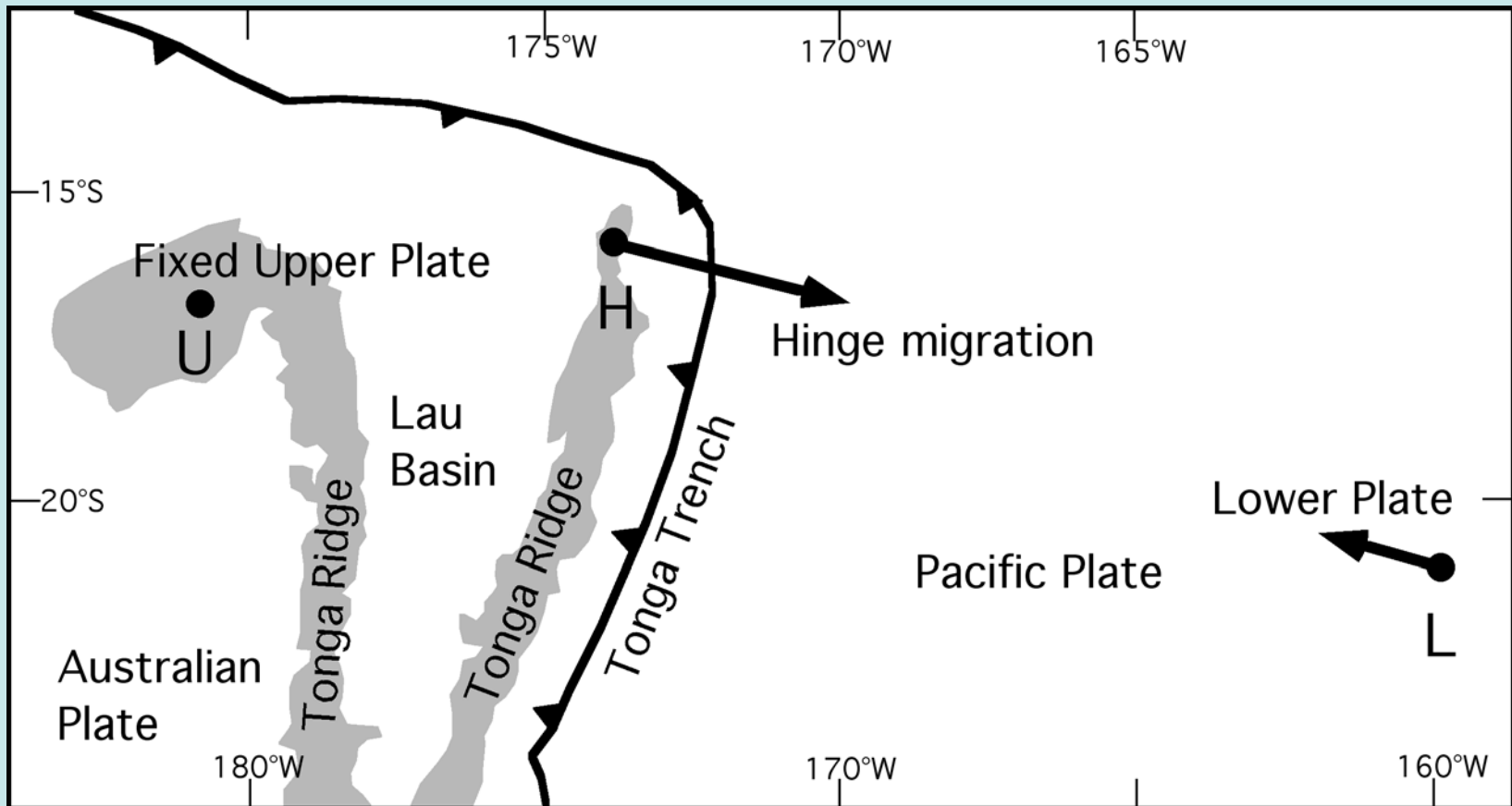


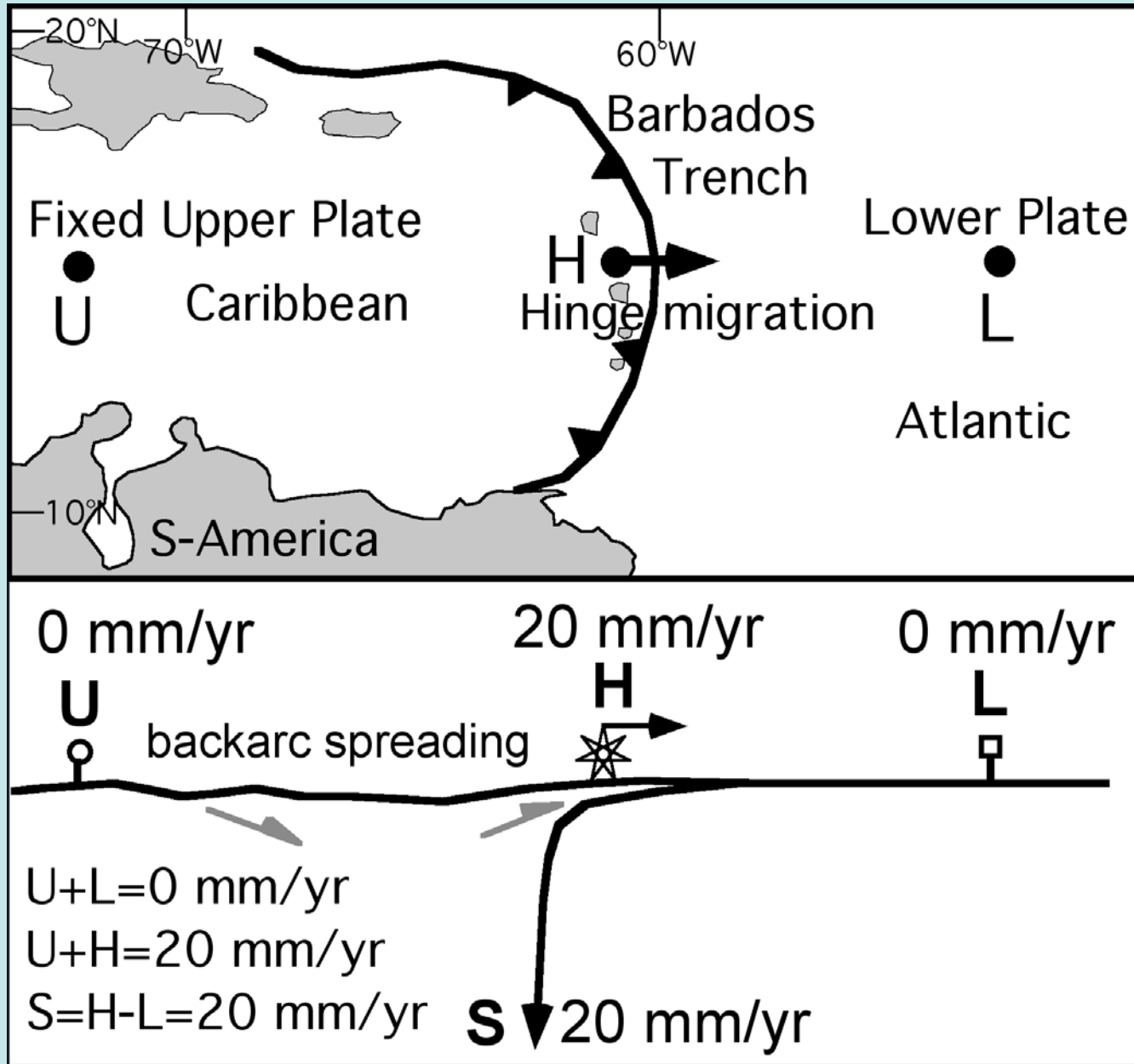
Subduction > Convergence

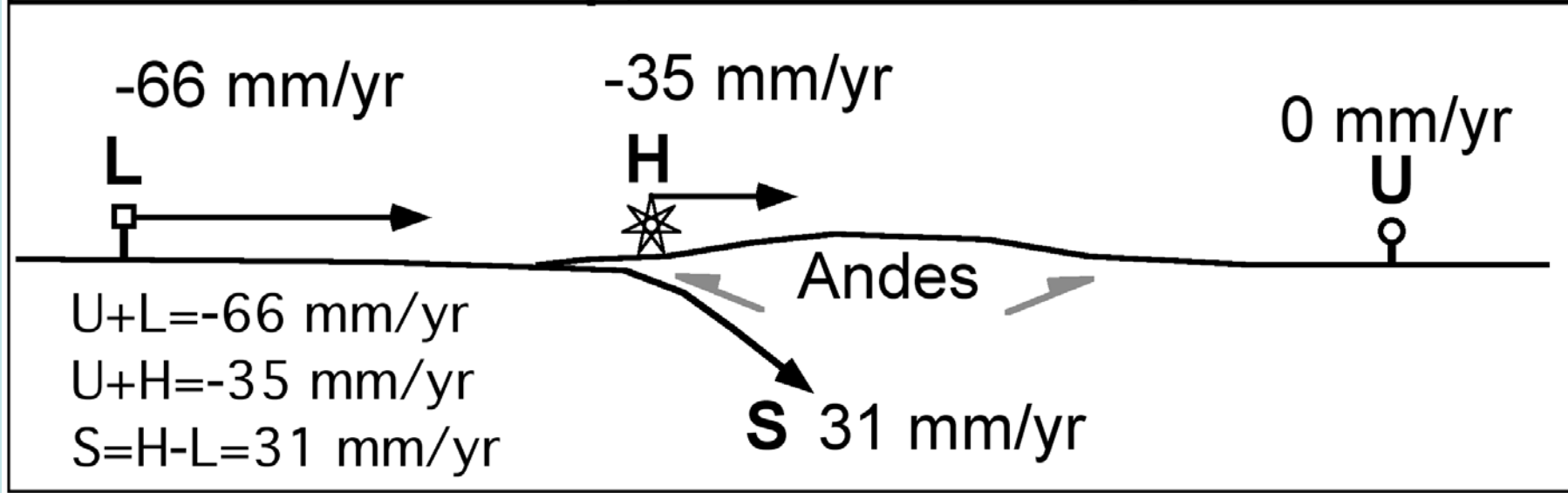
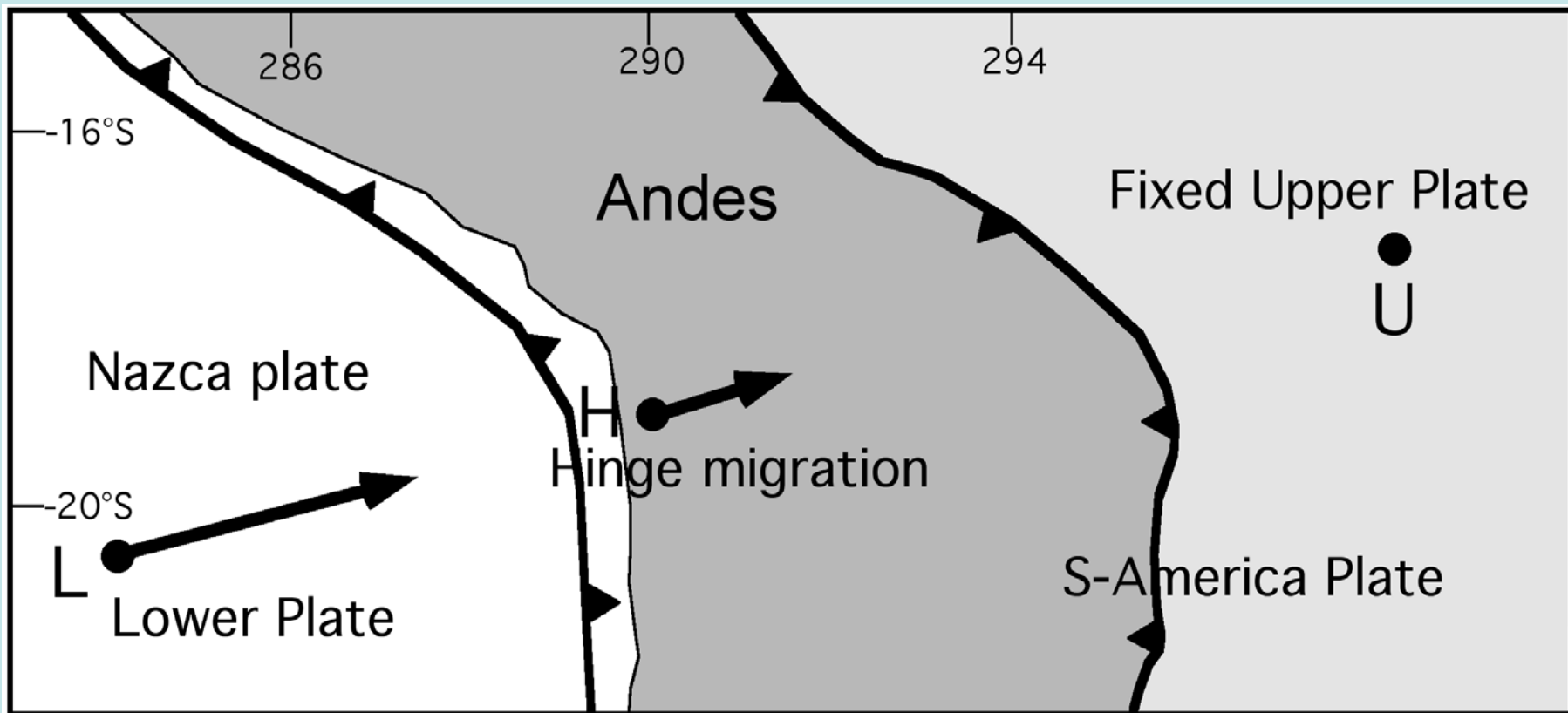
Apenninic type

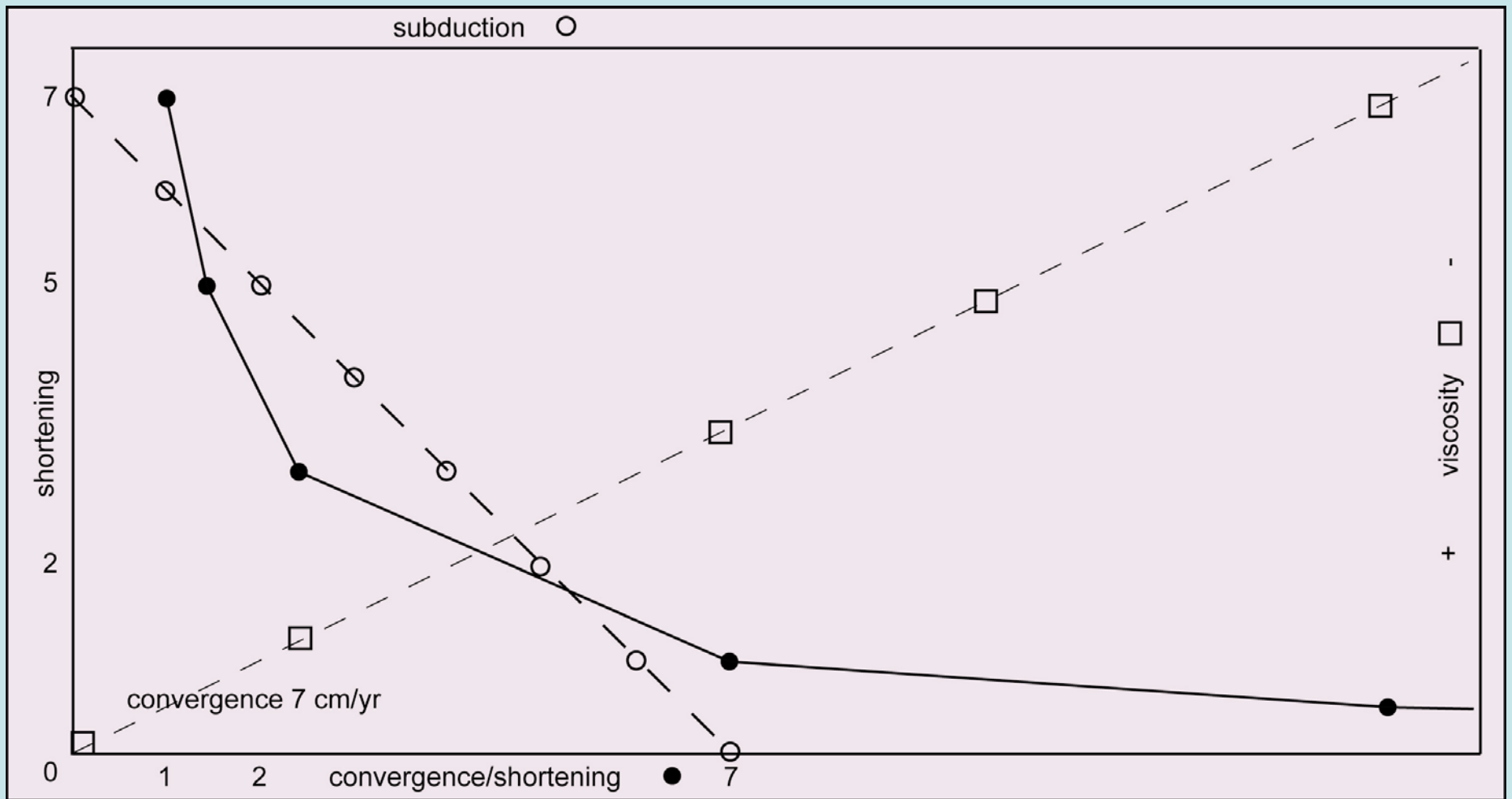
Alpine type

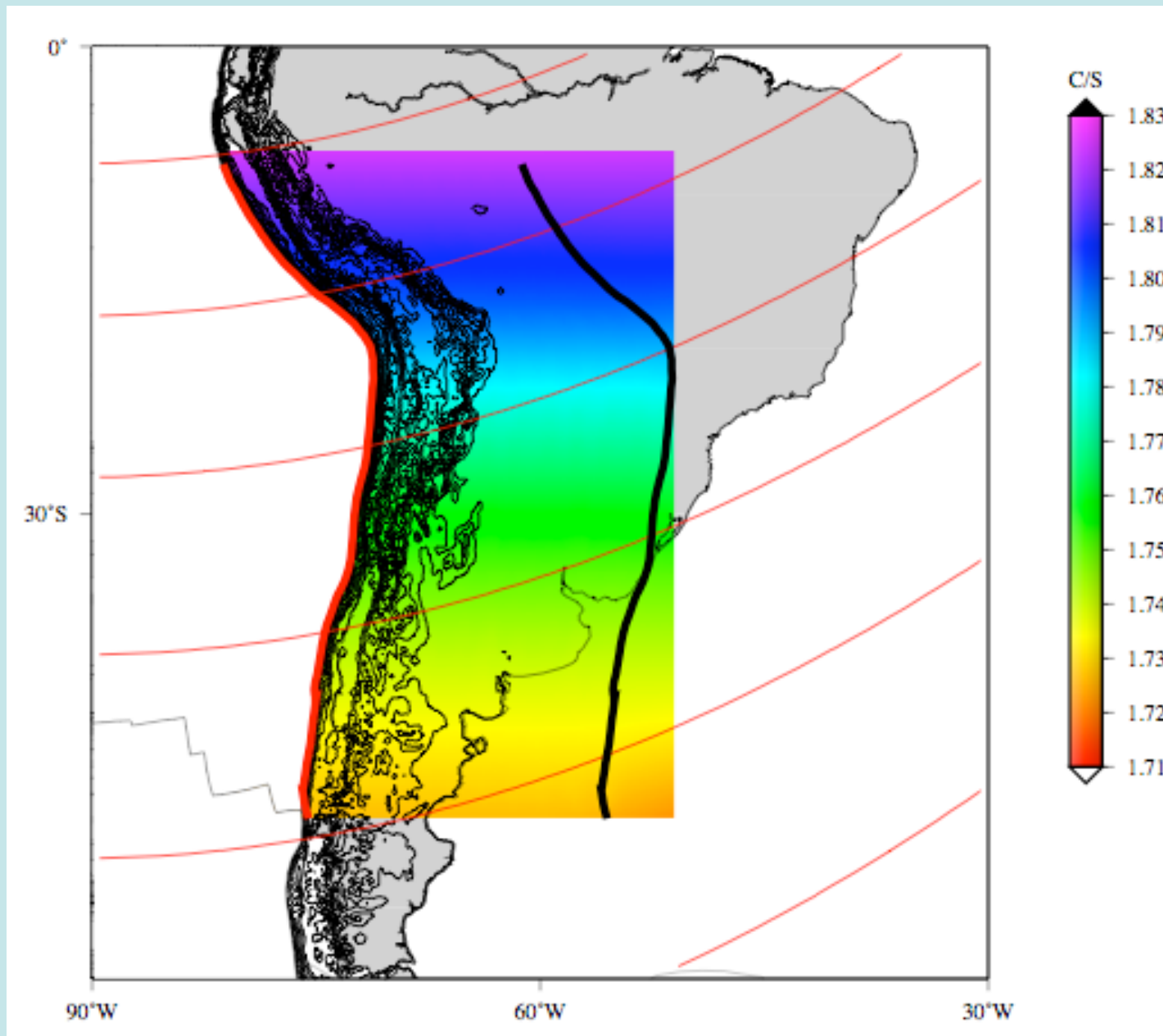




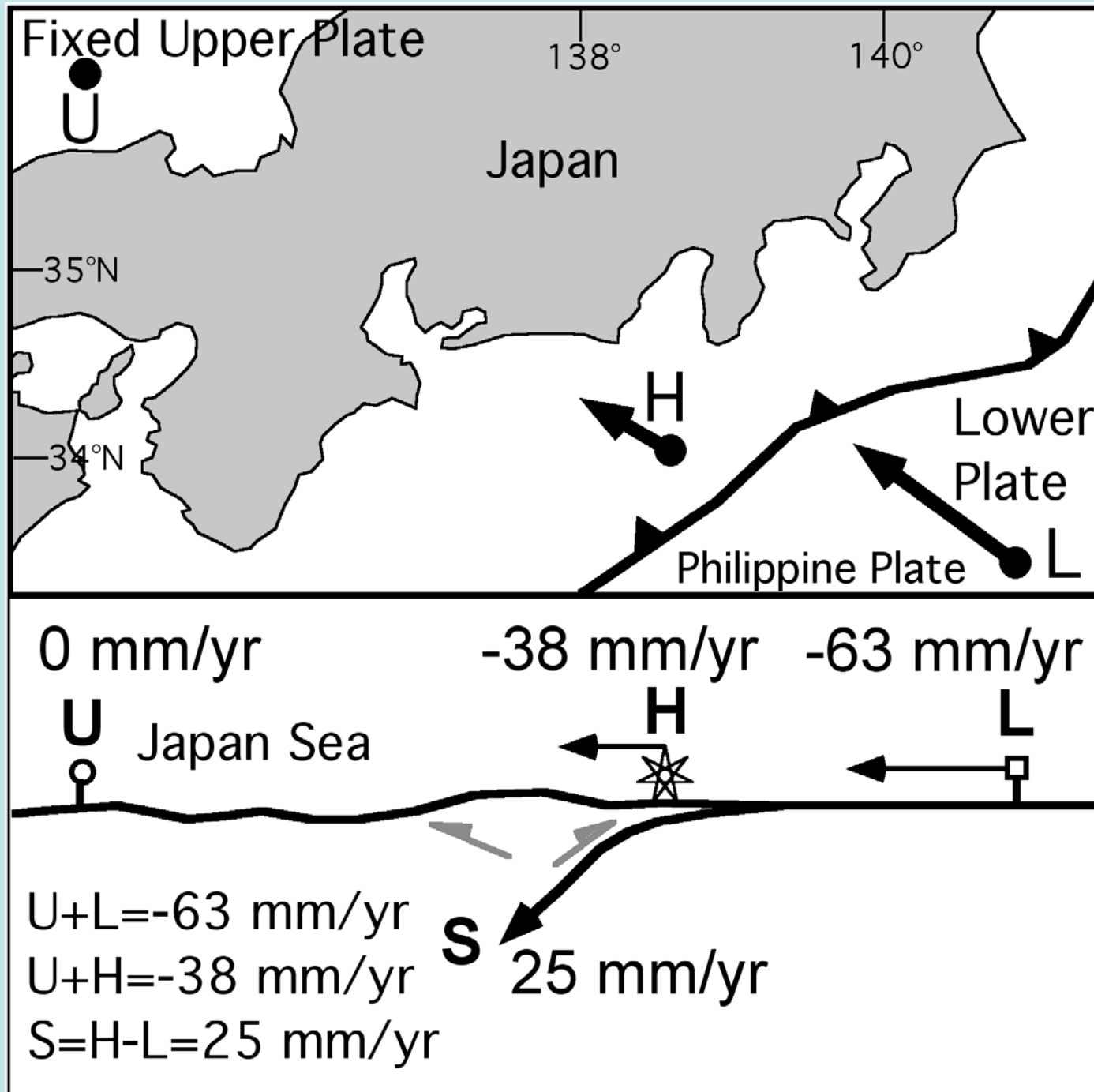




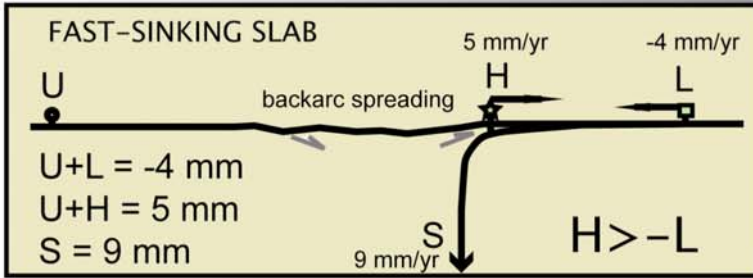




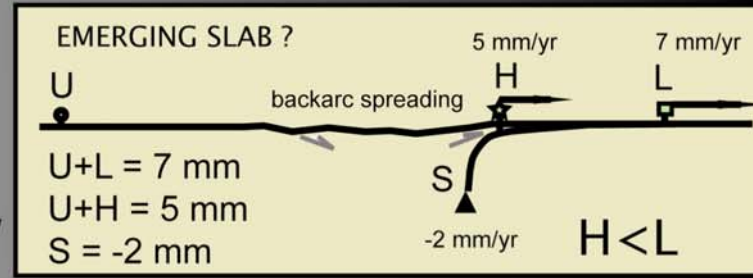
Higher convergence/shortening ratio, higher lithosphere viscosity



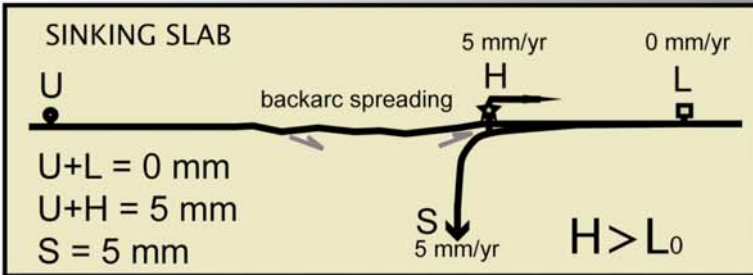
WEST



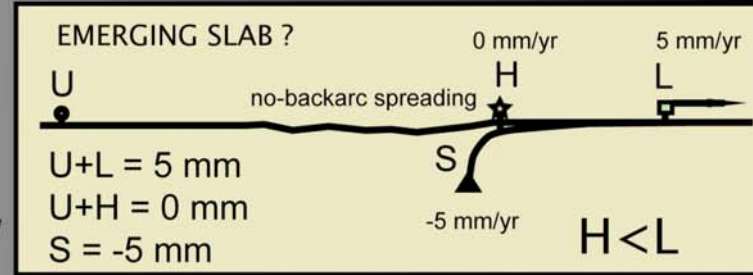
1W



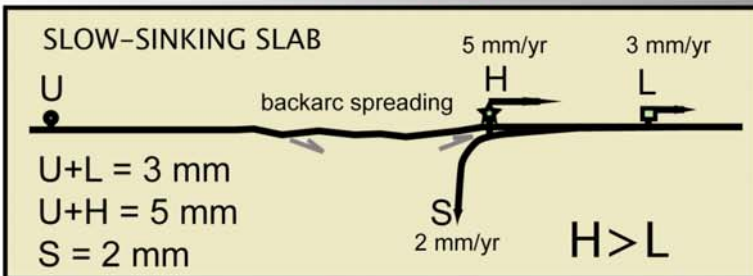
5W



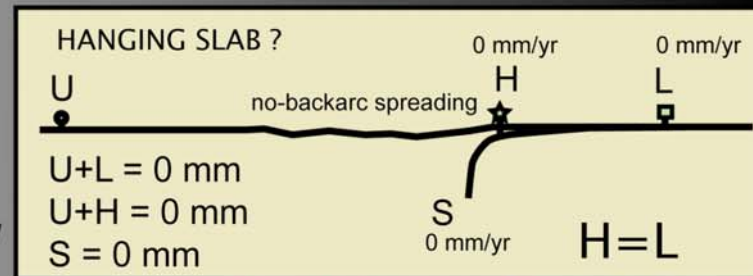
2W



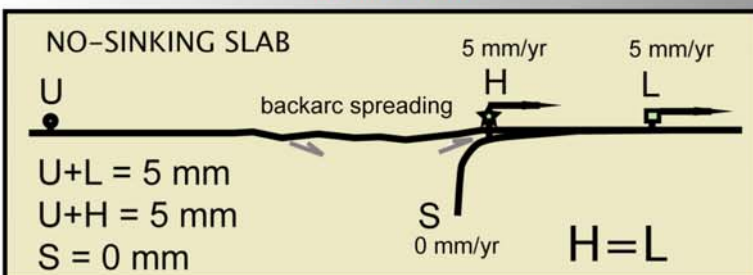
6W



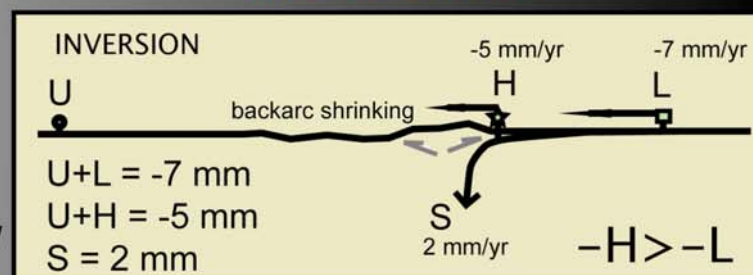
3W



7W

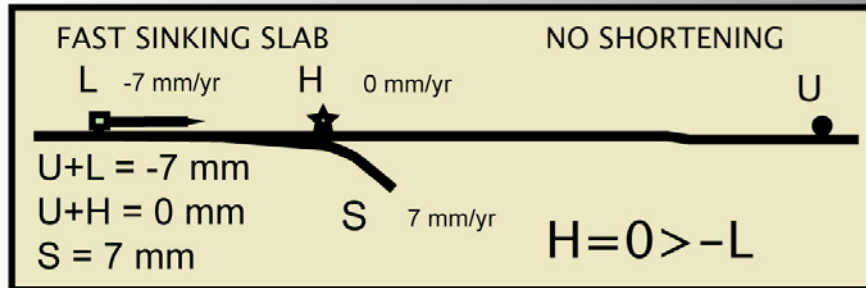


4W

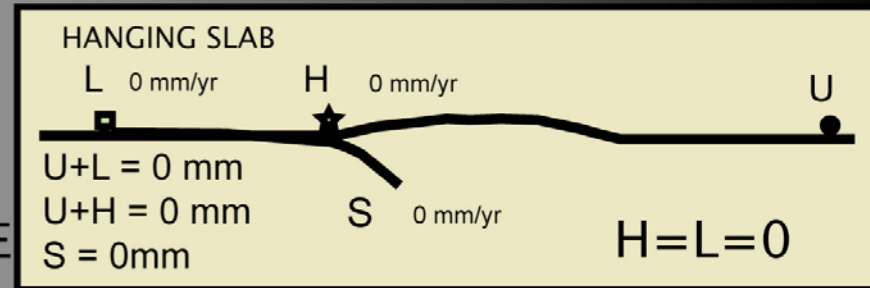


8W

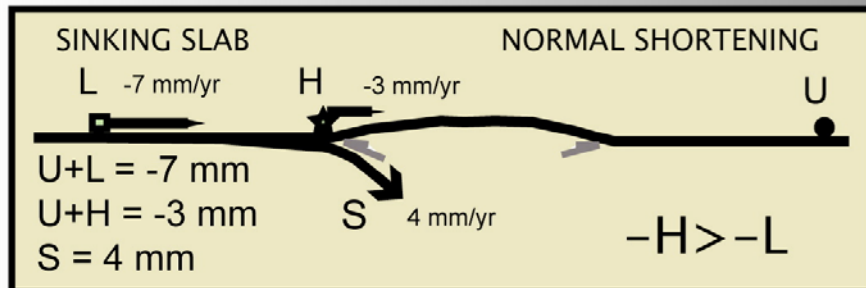
EAST-NORTHEAST



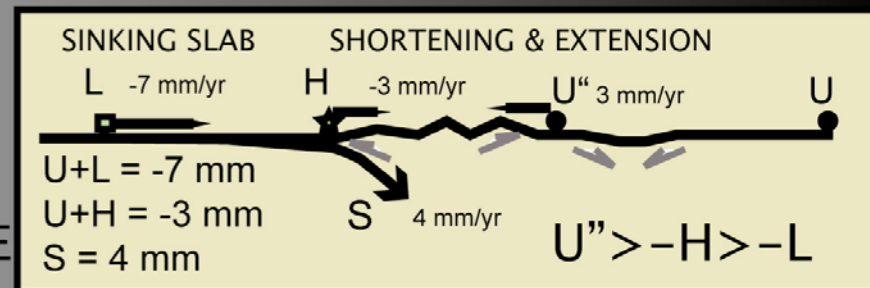
1E



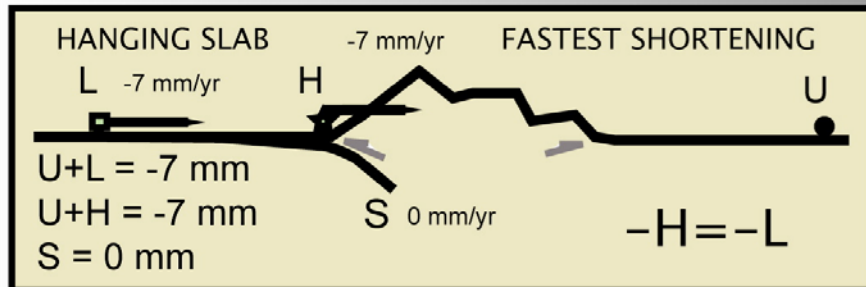
4E



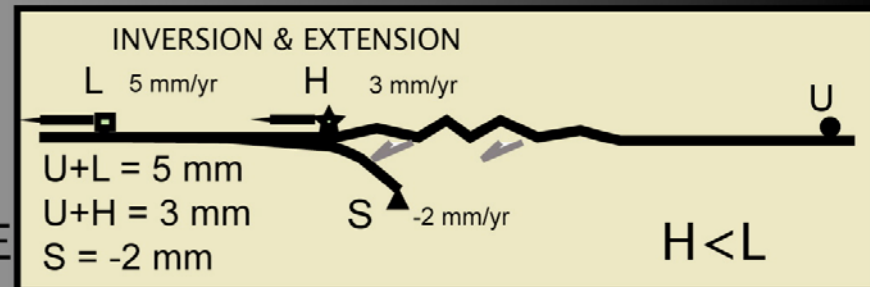
2E



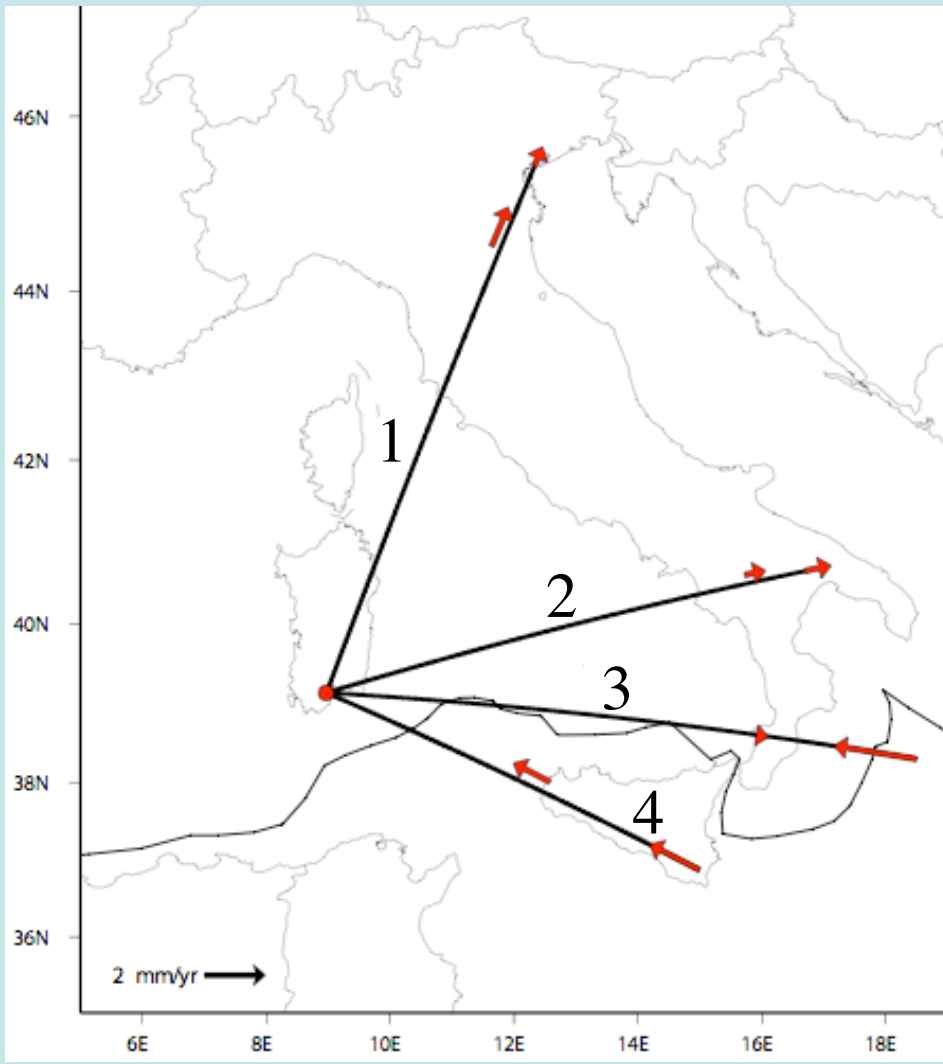
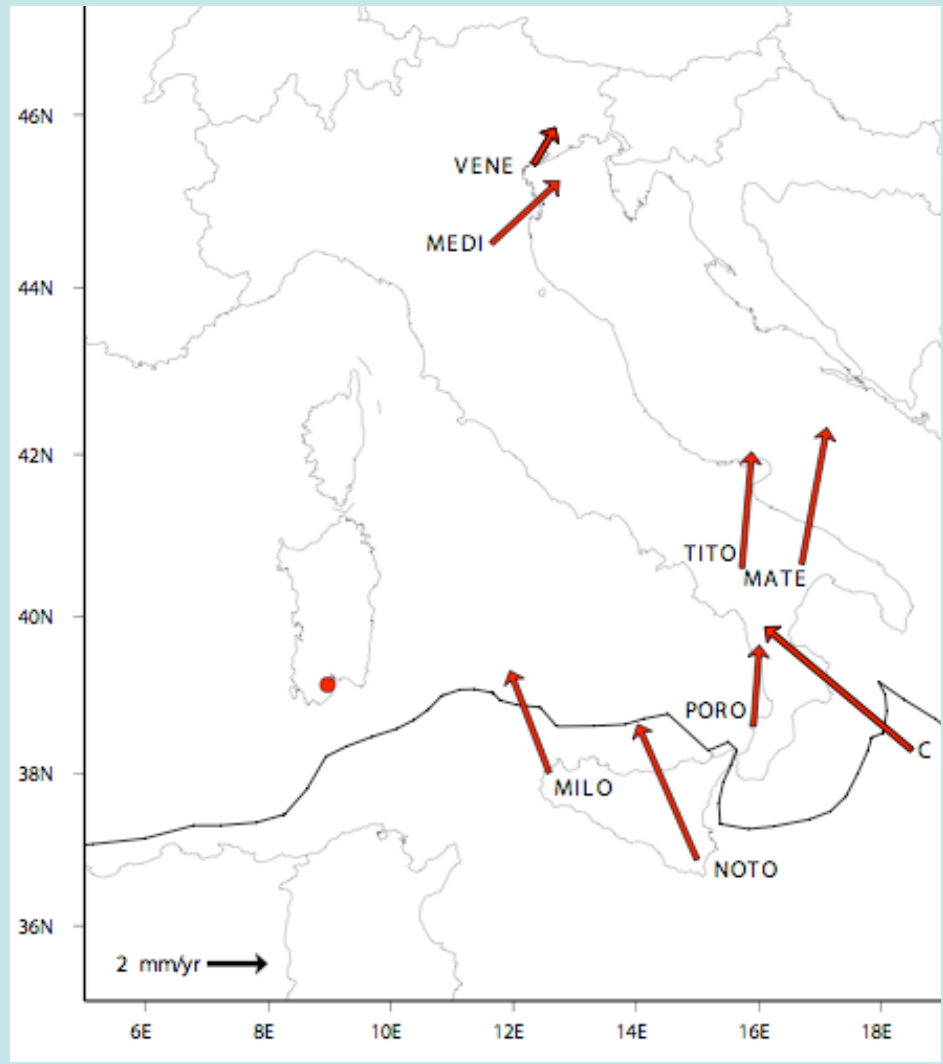
5E



3E



6E



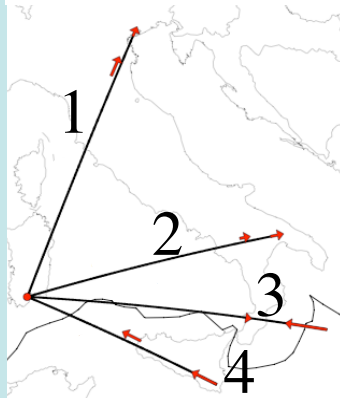
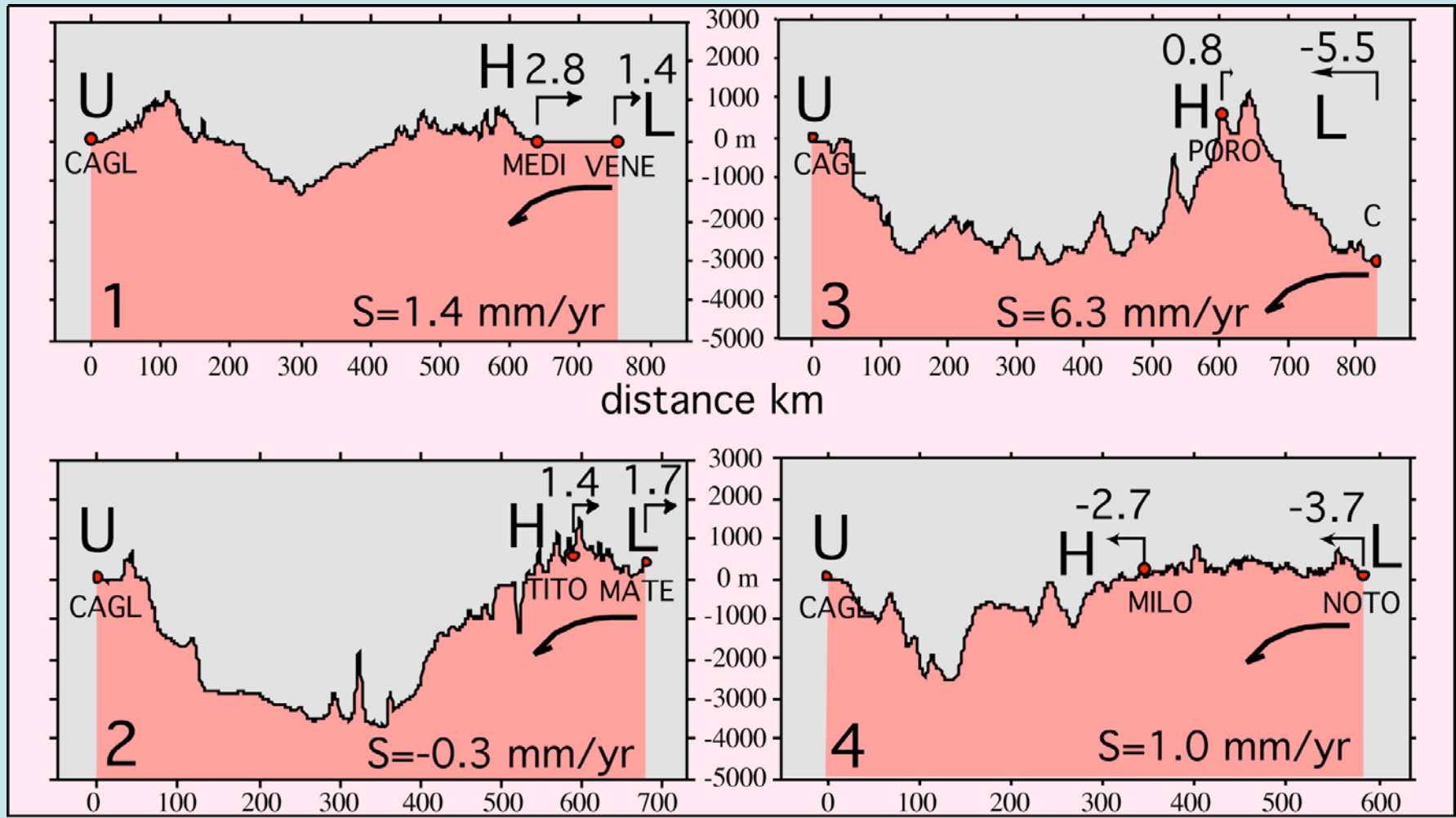
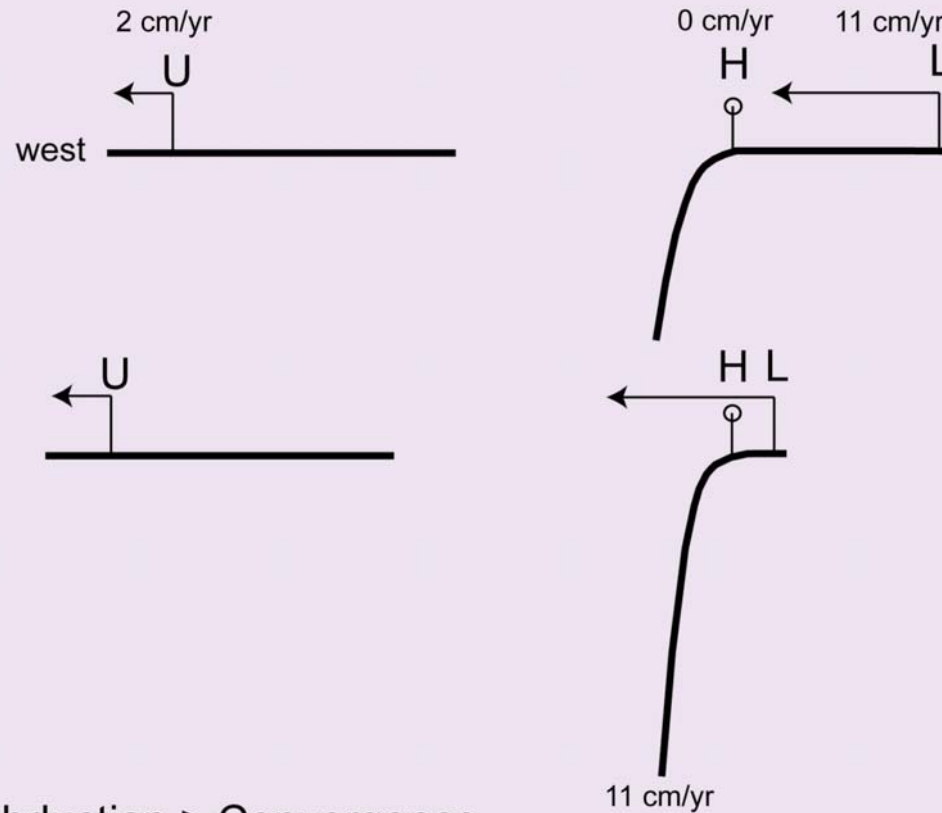


PLATE MOTIONS RELATIVE TO THE MANTLE

L = Lower plate
H = Subduction hinge
U = Upper plate

Convergence 9 cm/yr
Shortening in the prism 11 cm/yr
Subduction 11 cm/yr
Backarc spreading 2 cm/yr



Subduction > Convergence
Hinge fixed with the mantle
Hinge diverging east relative to upper plate

PLATE MOTIONS RELATIVE TO THE MANTLE

L = Foreland plate
H = Subduction hinge
U = Upper plate

Convergence 7 cm/yr
Shortening in the orogen 5 cm/yr
Subduction 2 cm/yr

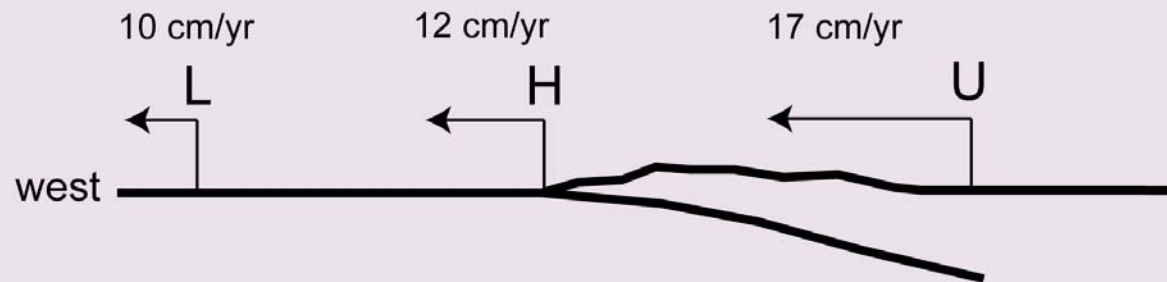
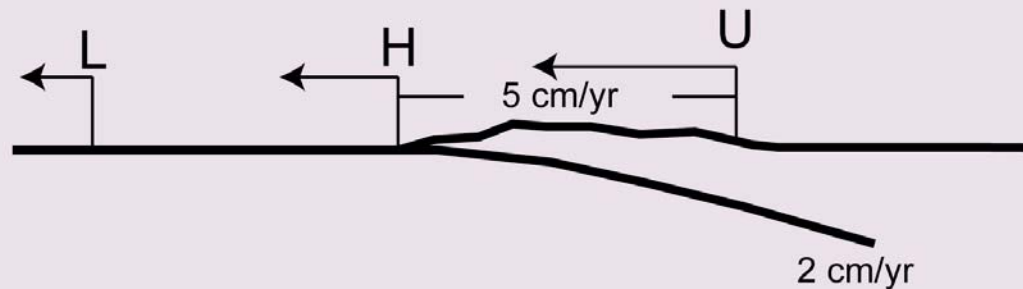
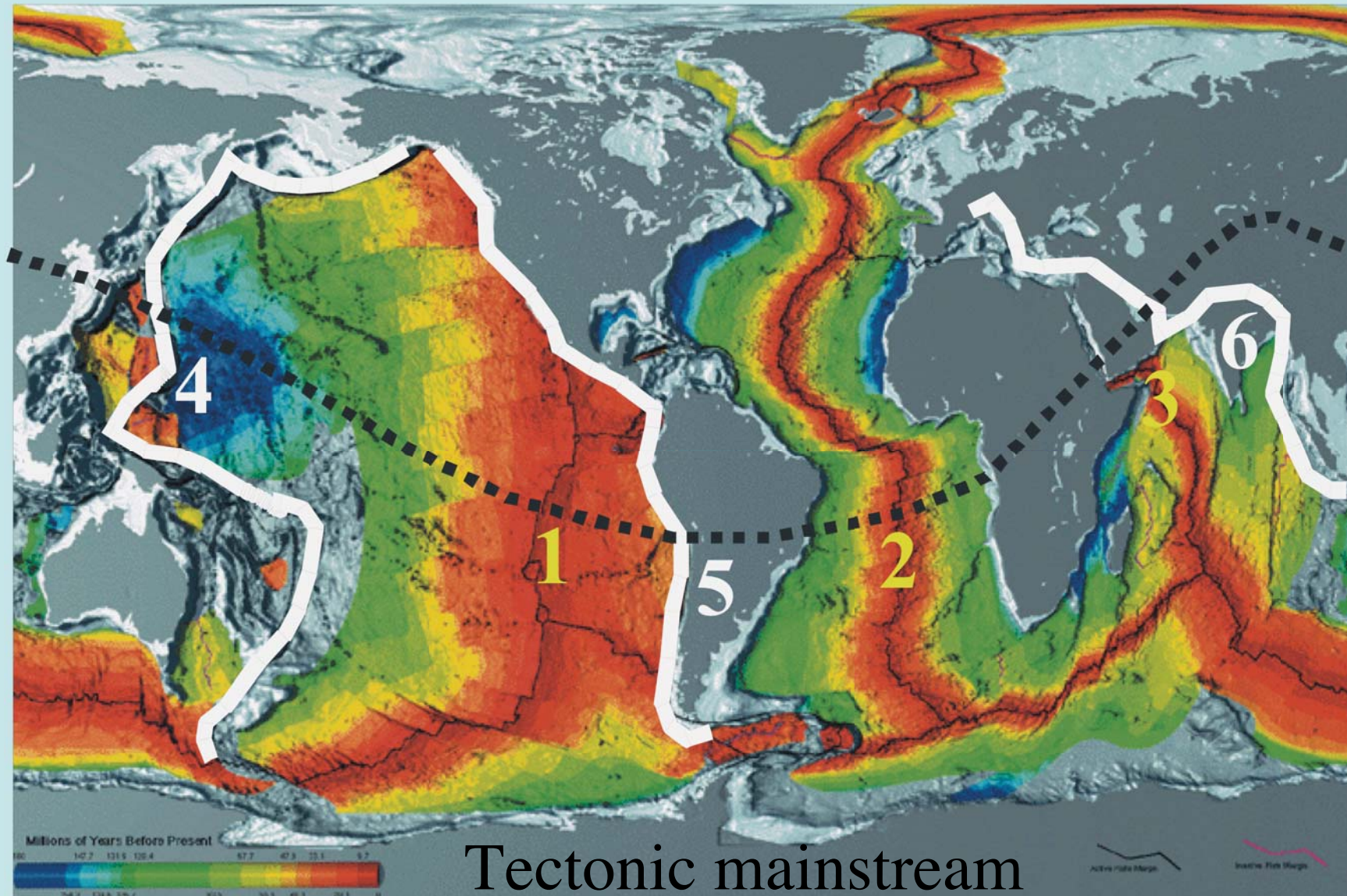


Plate motions not controlled by subduction rate



Subduction < Convergence
Hinge retreating west relative to mantle
Hinge converging east relative to upper plate



1= East Pacific Rise

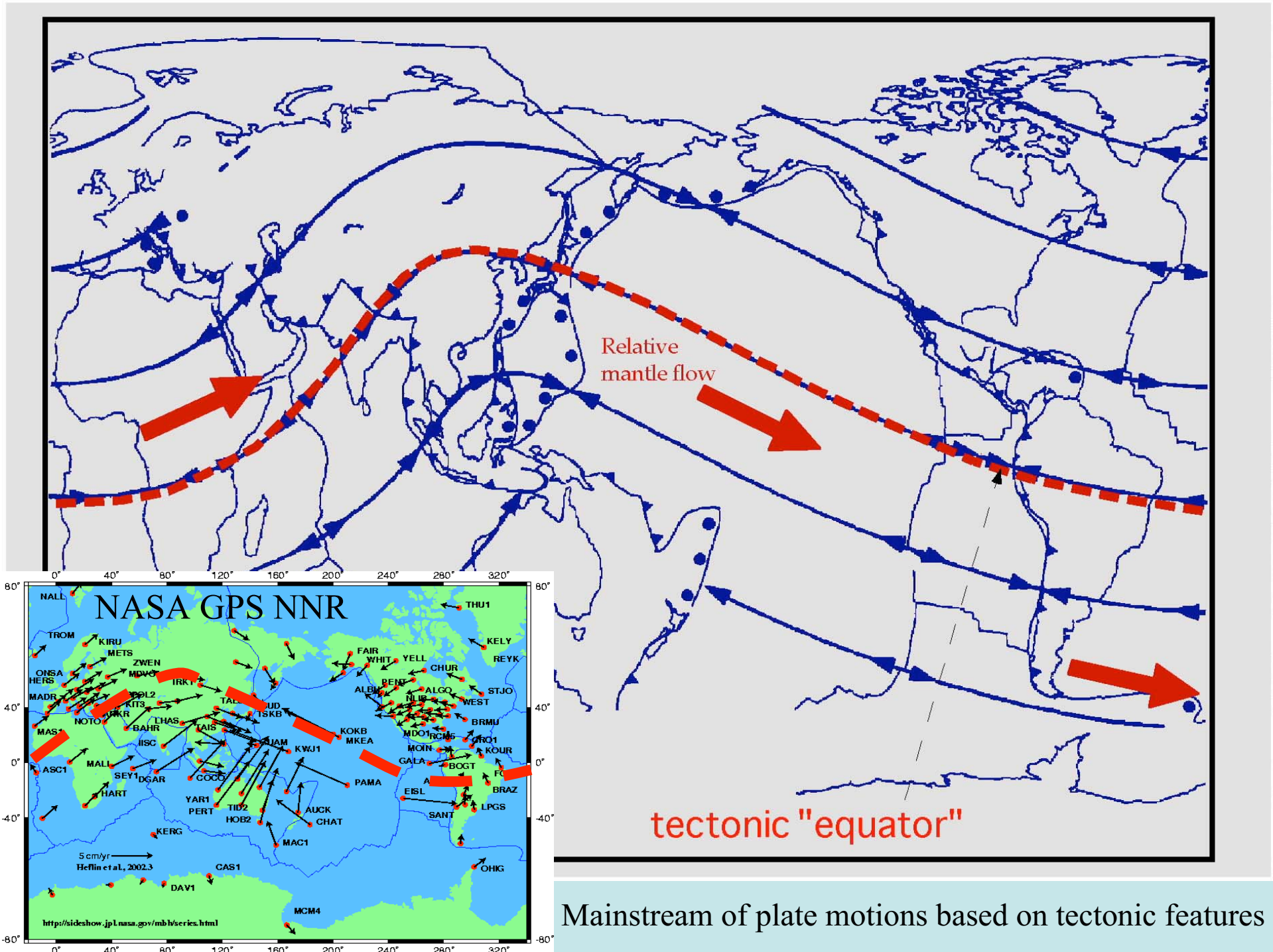
2= Atlantic Rift

3= Red Sea Indian Ocean Rift

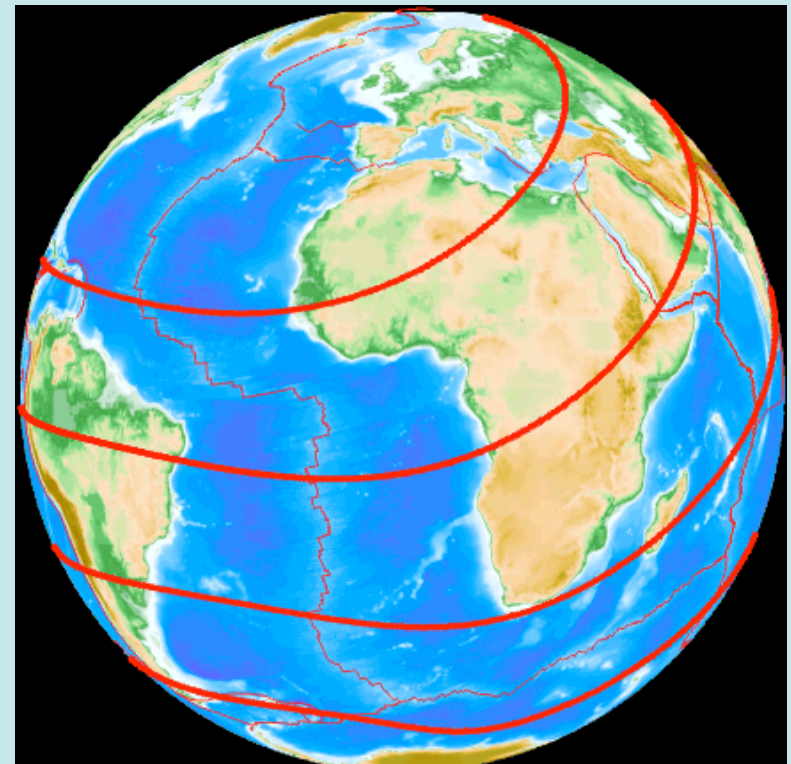
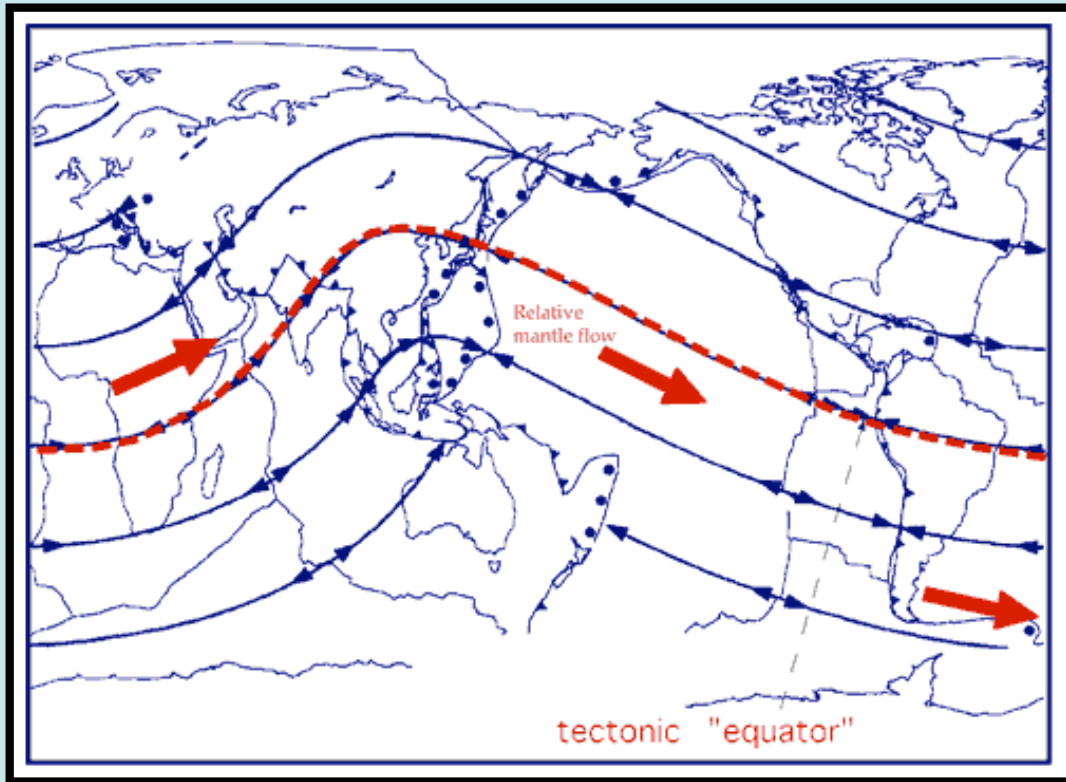
4= West Pacific Subduction

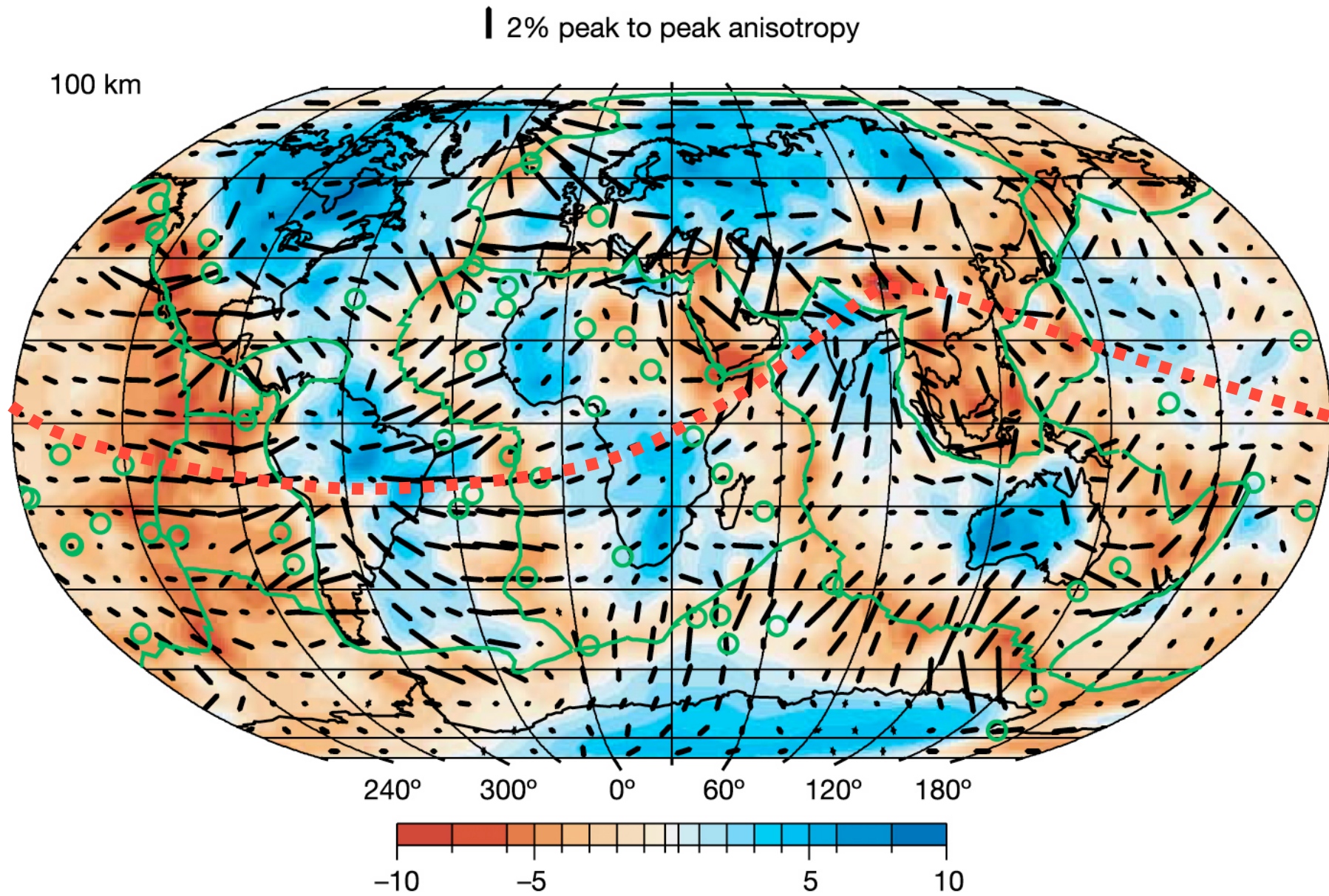
5= Andean Subduction

6= Zagros-Himalayas Subduction



Mainstream of plate motions based on tectonic features





Debayle et al., 2005, Nature

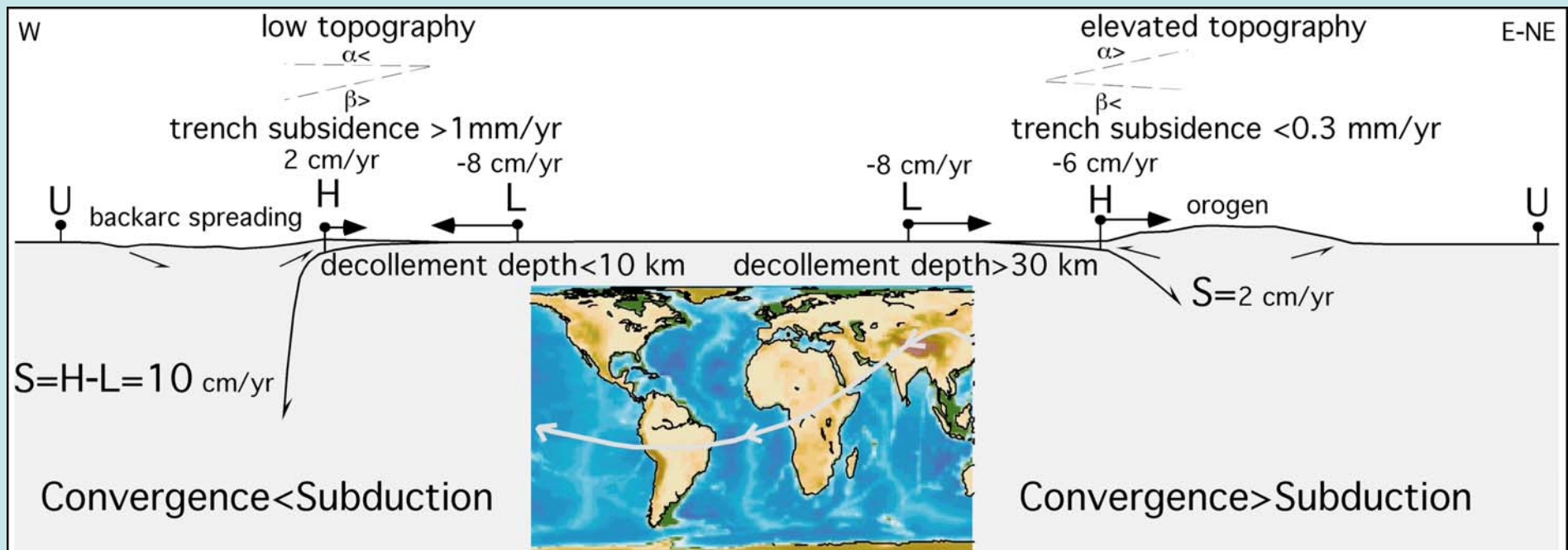
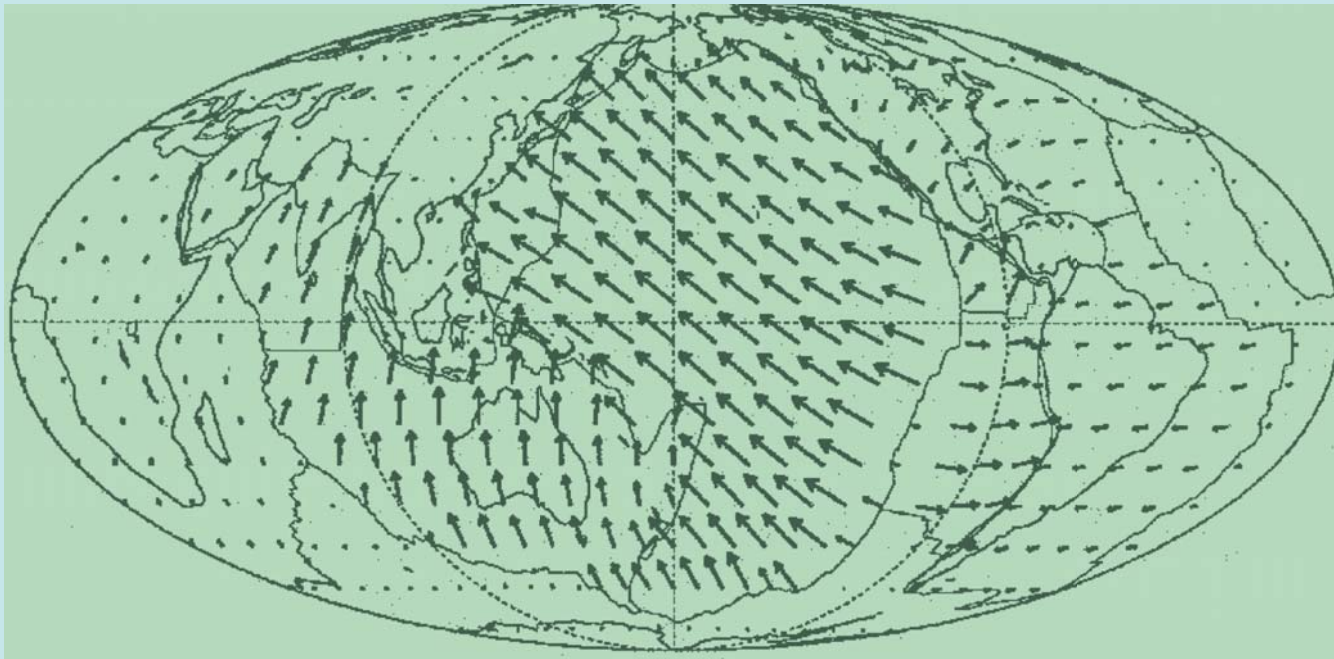
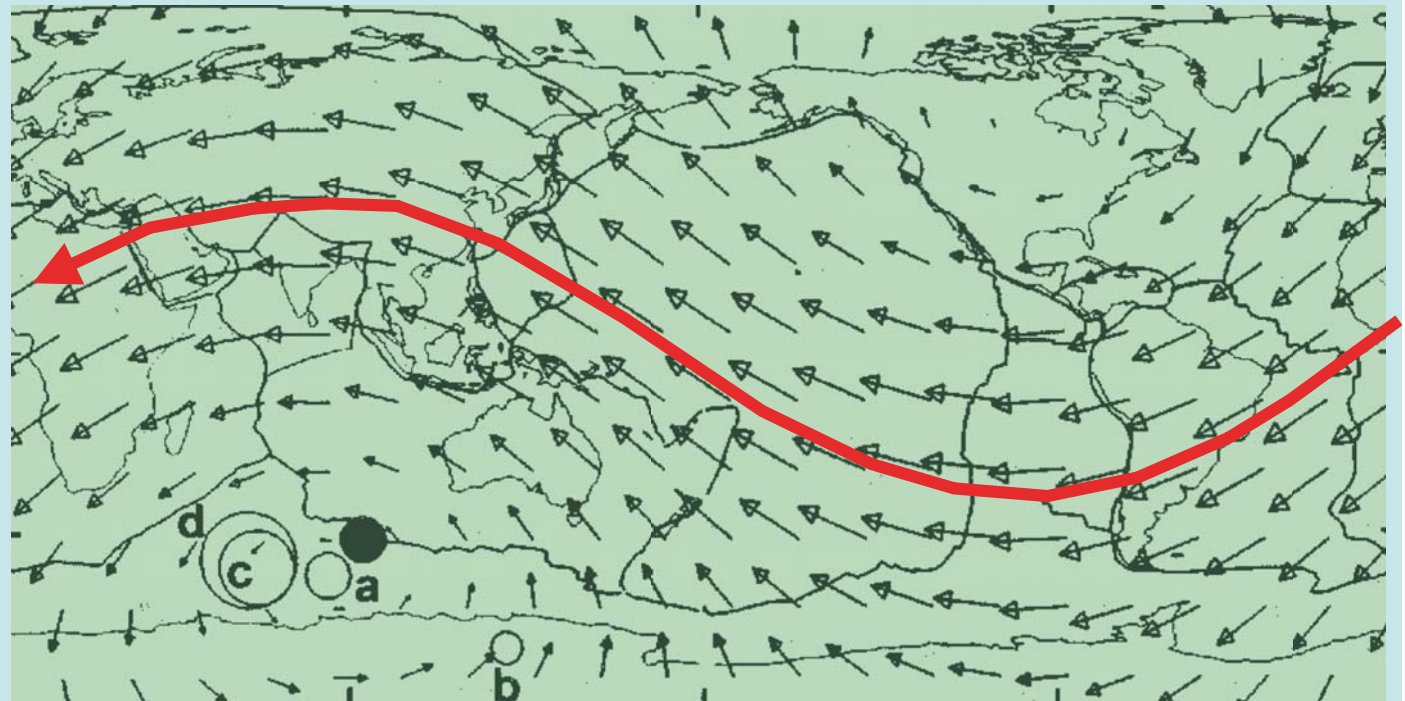
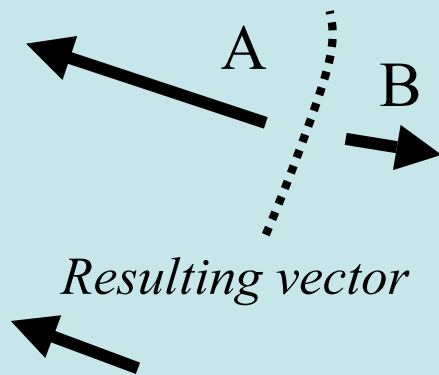


Plate motions based on the
Hotspot Reference Frame
(O'Connell, Gable & Hager, 1991)

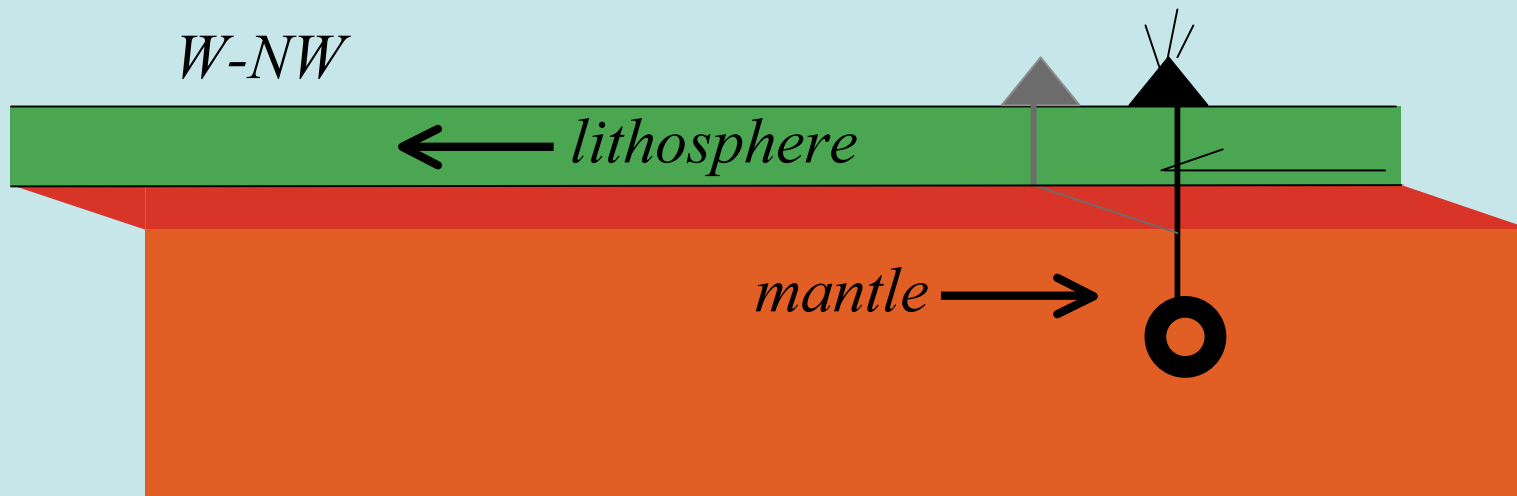


(Ricard, Doglioni & Sabadini, JGR, 1991)

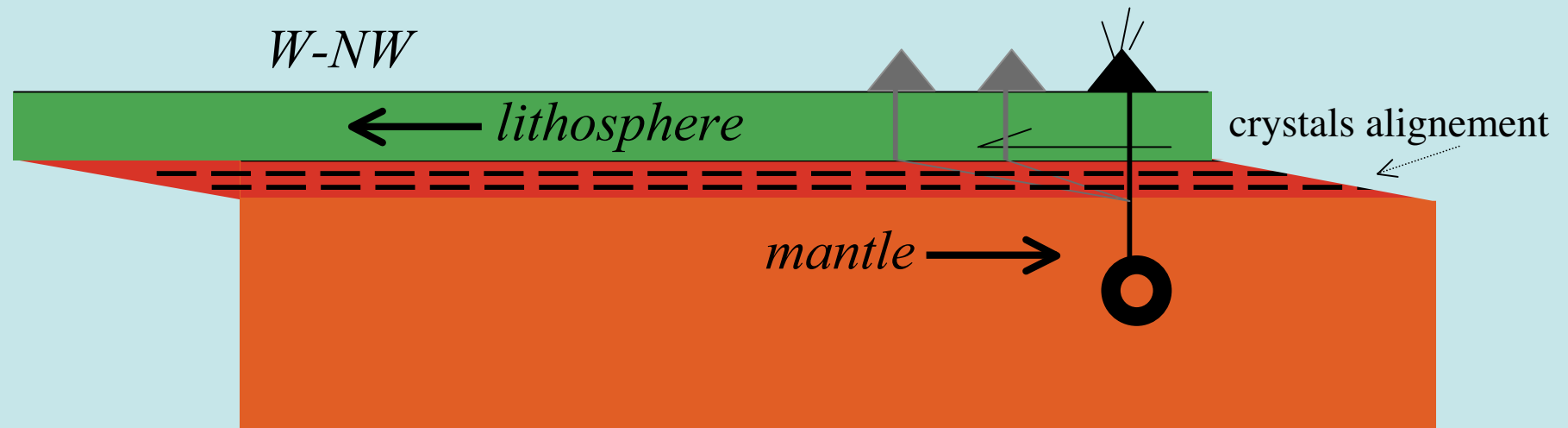


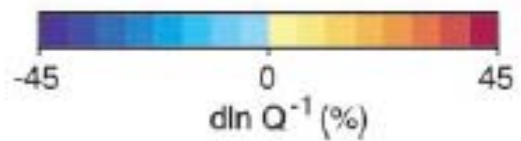
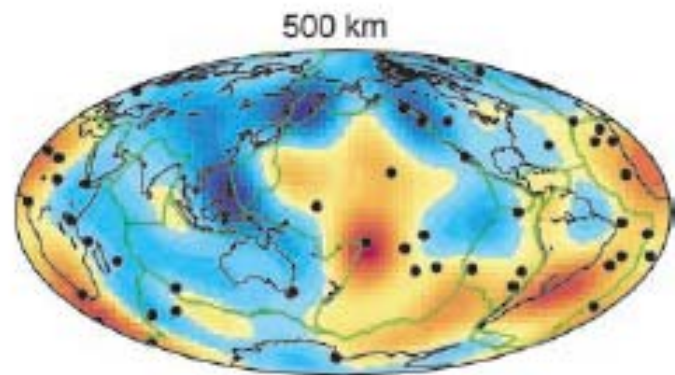
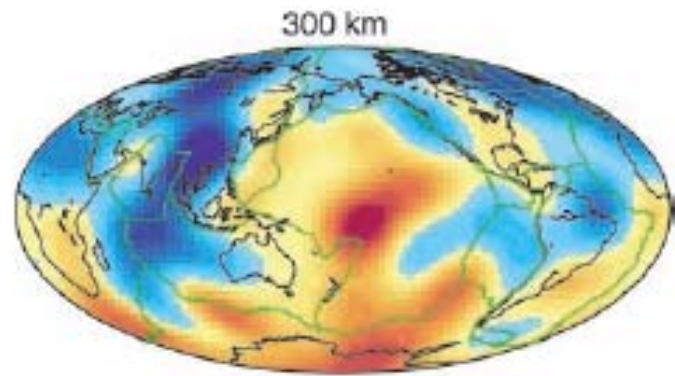


“Westward” drift of the lithosphere

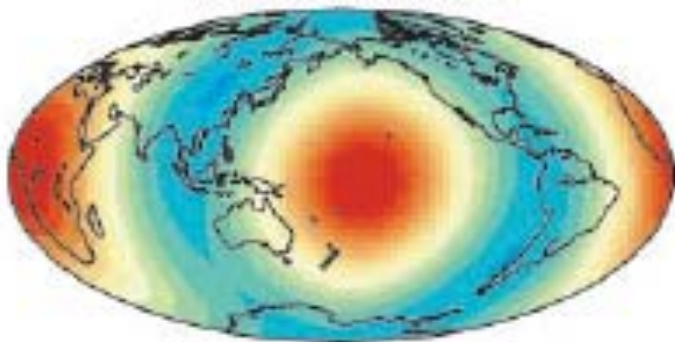


“Westward” drift of the lithosphere (or “eastward” mantle flow)

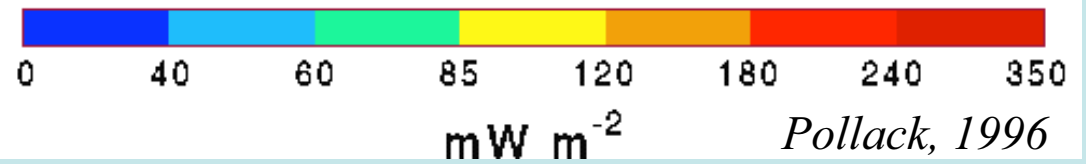
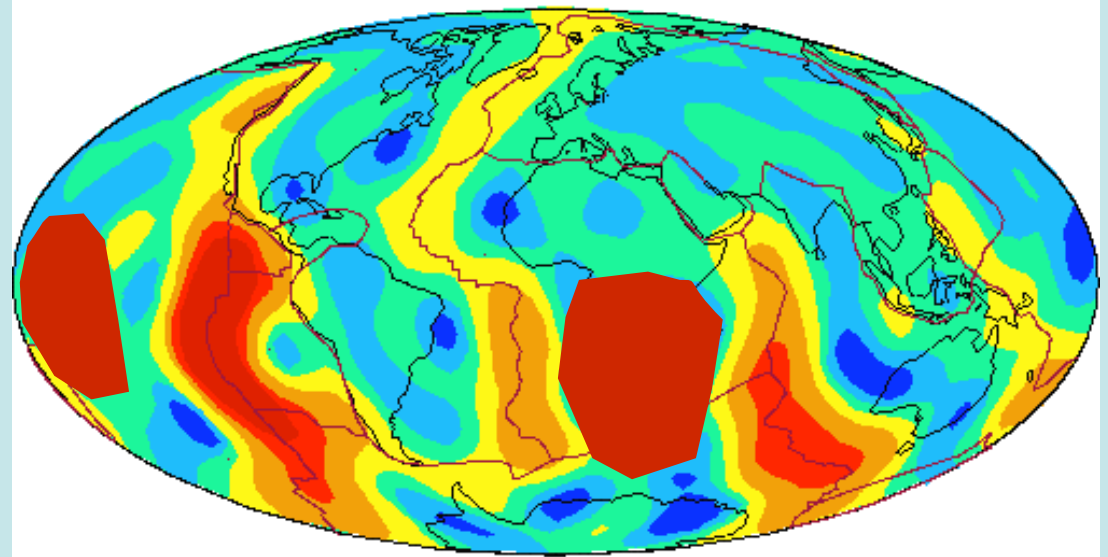




VSH at 2800 km

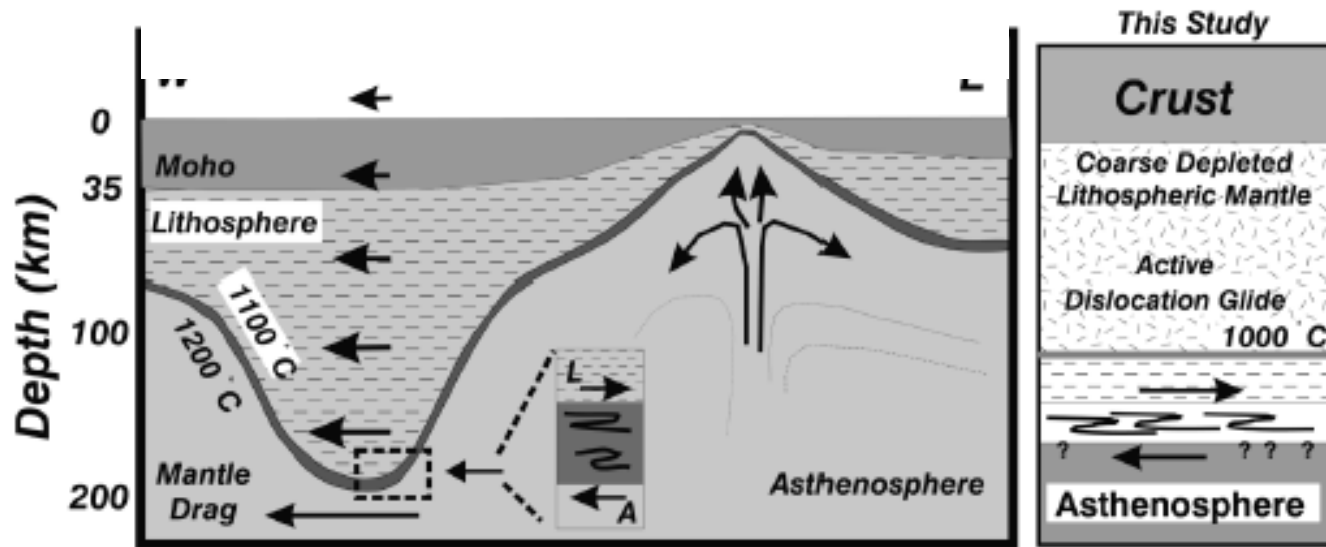
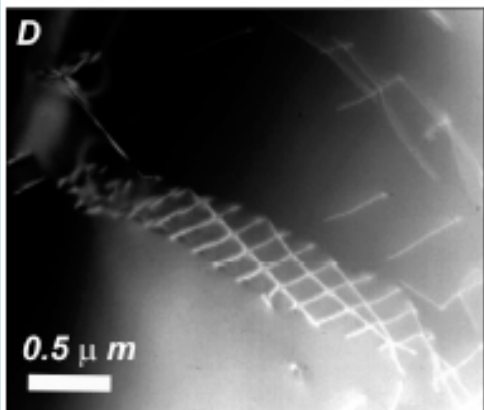
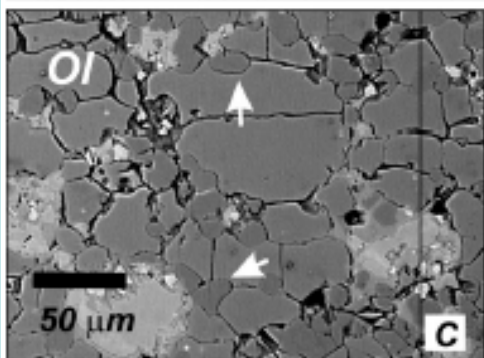
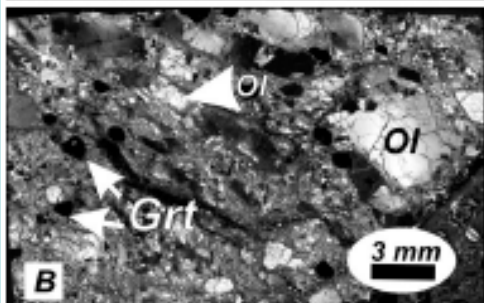
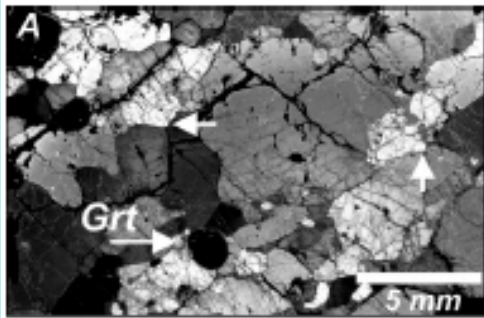


Heat Flow

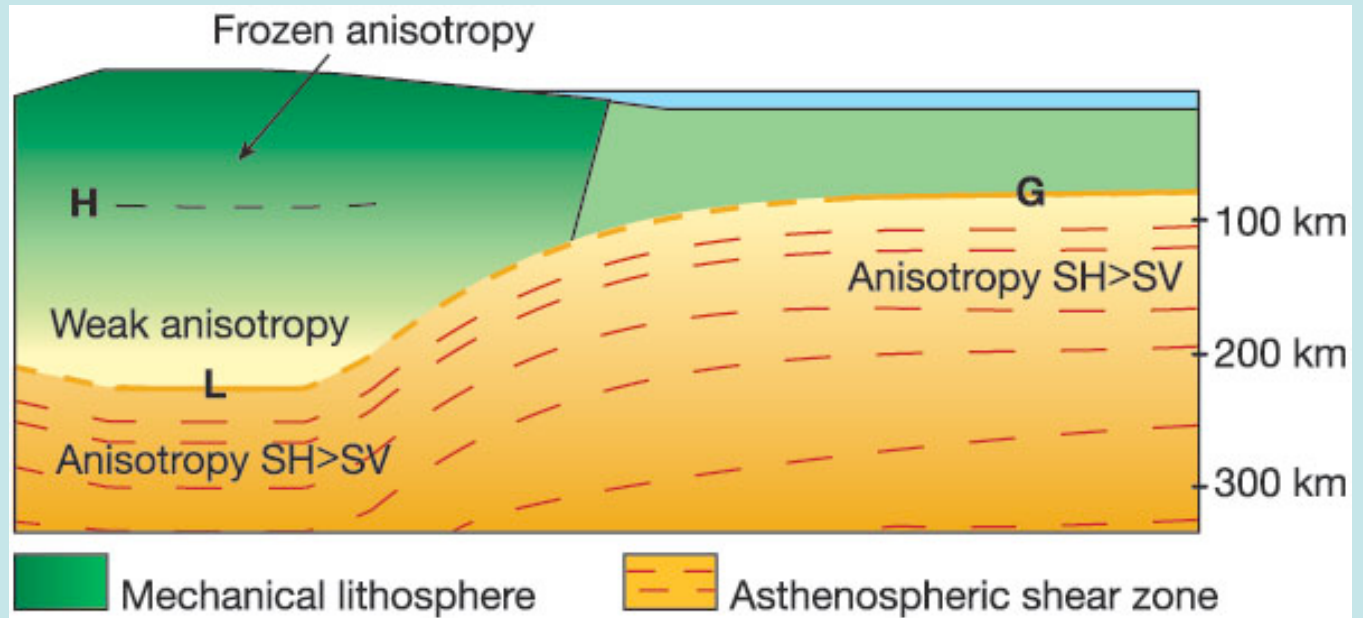


Pollack, 1996

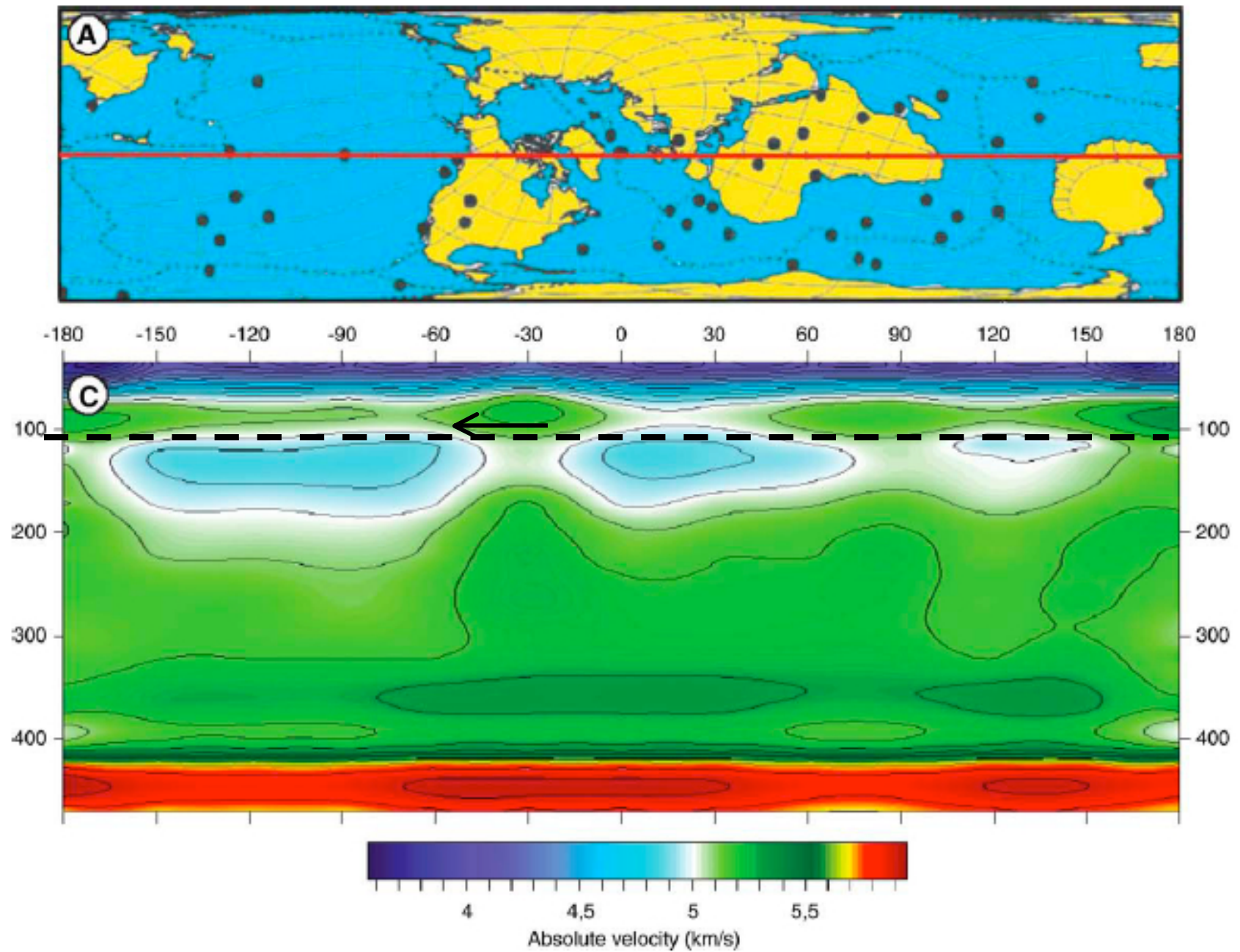
Romanowicz & Gung, 2002



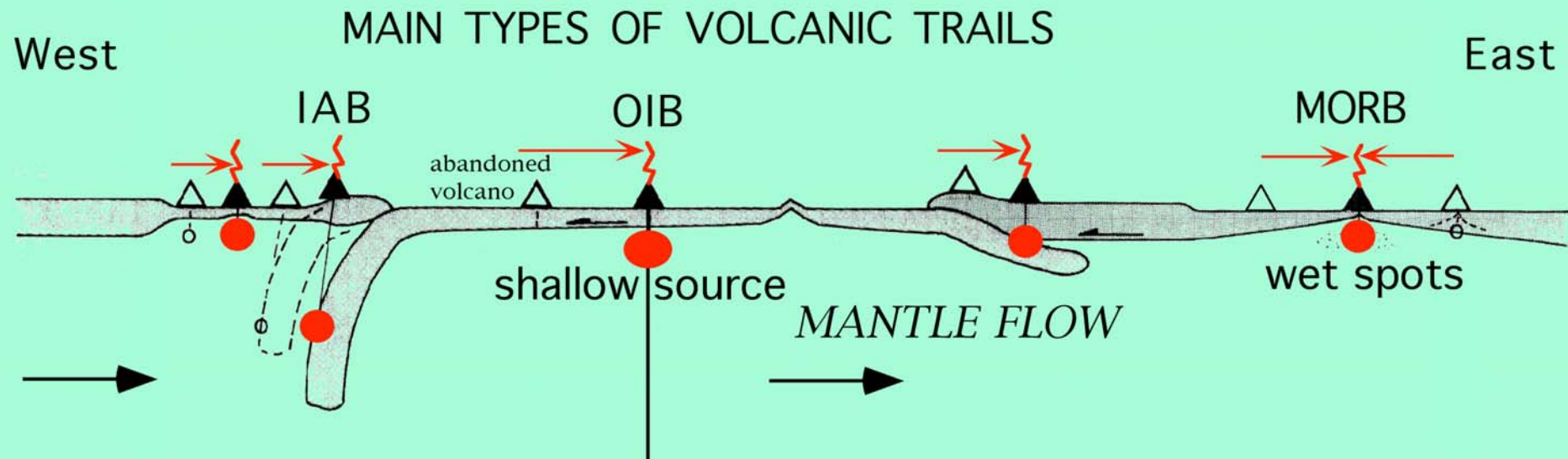
Kennedy et al., Geology 2002



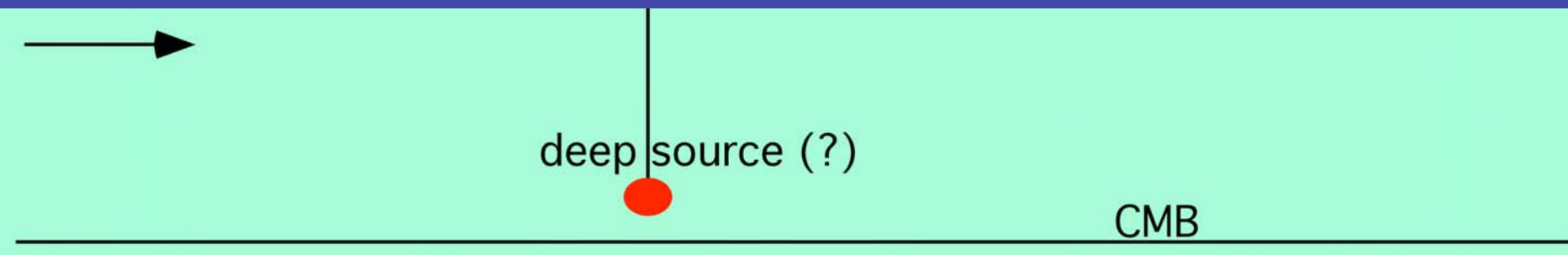
Gung et al., Nature 2003



- Different depths of the magmatic sources
- Most of them in the asthenosphere or in the lithospheric mantle
- Most of them along moving and mantle-detached plate margins



*Without "hotspots" located along plate margins & shallow plumes
Westward drift = $>100 \text{ mm} \cdot \text{yr}^{-1}$*

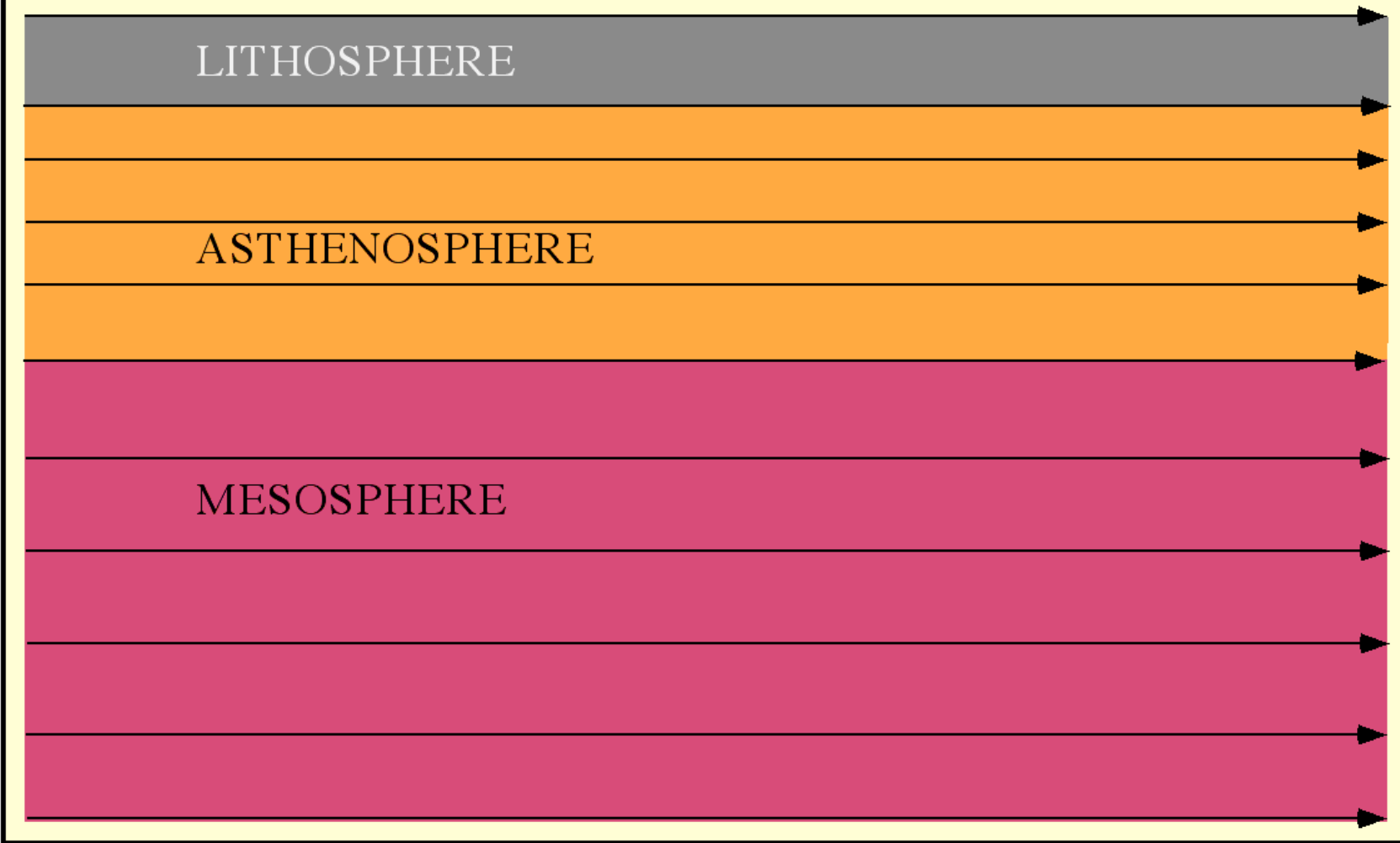


WEST

LITHOSPHERE

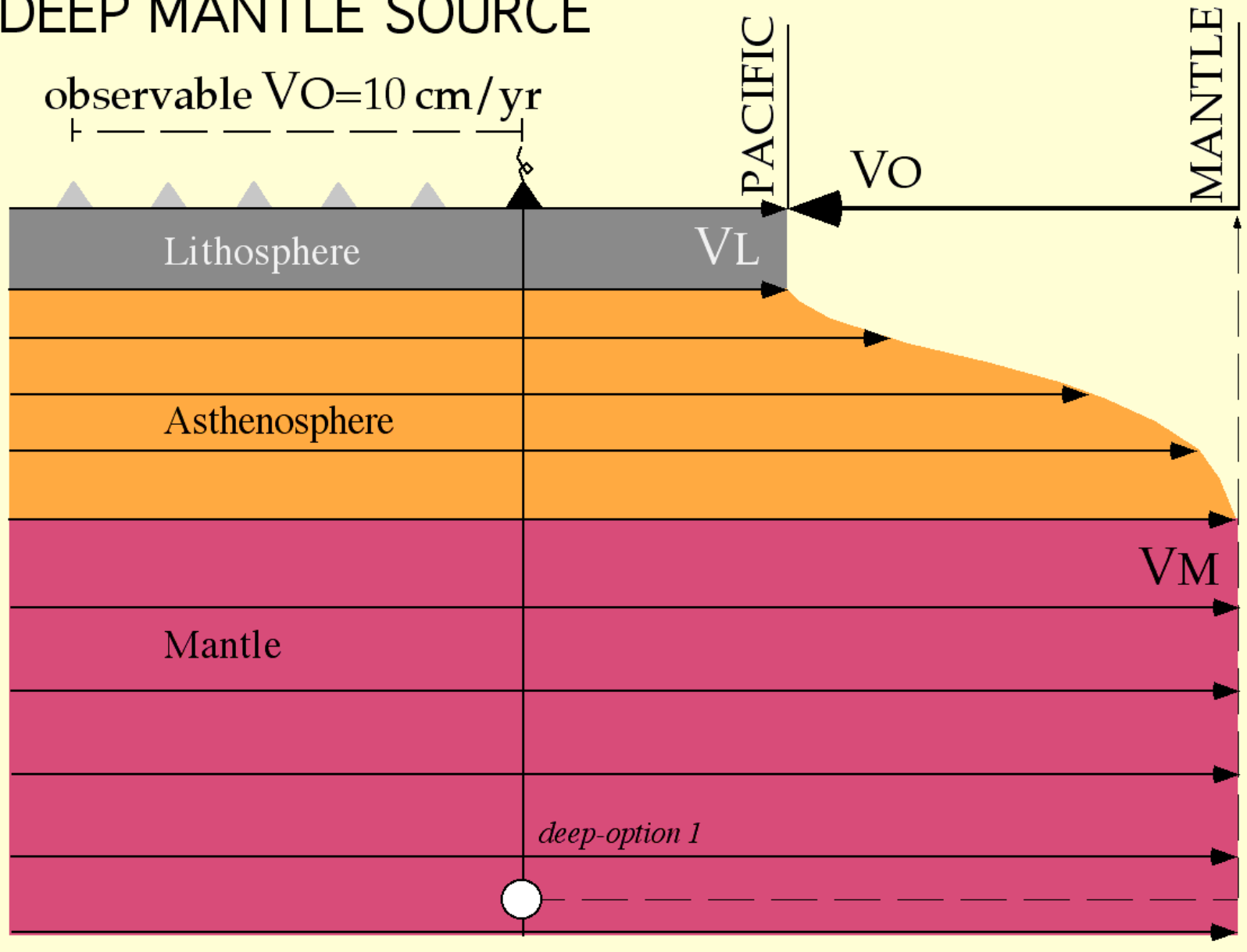
ASTHENOSPHERE

MESOSPHERE

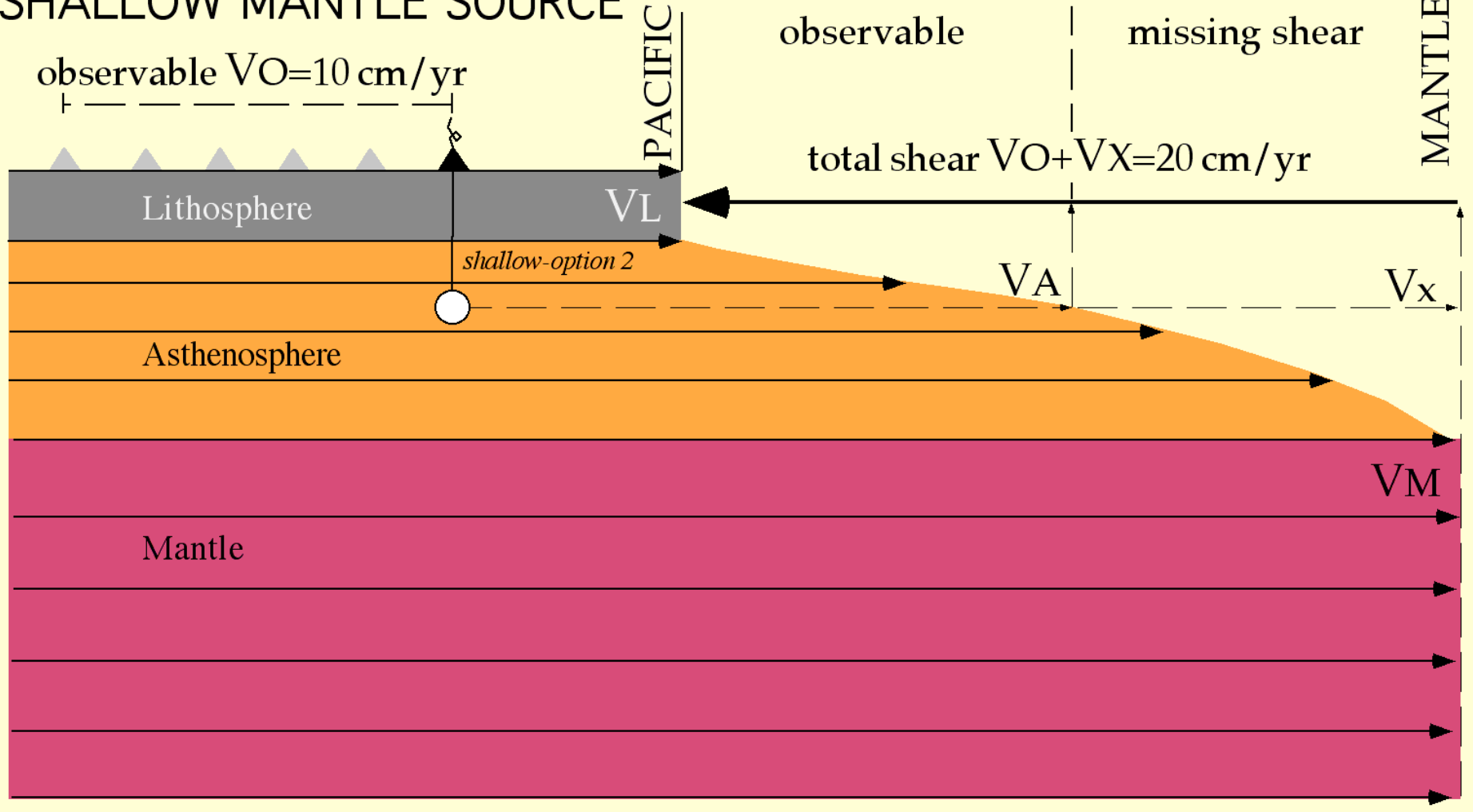


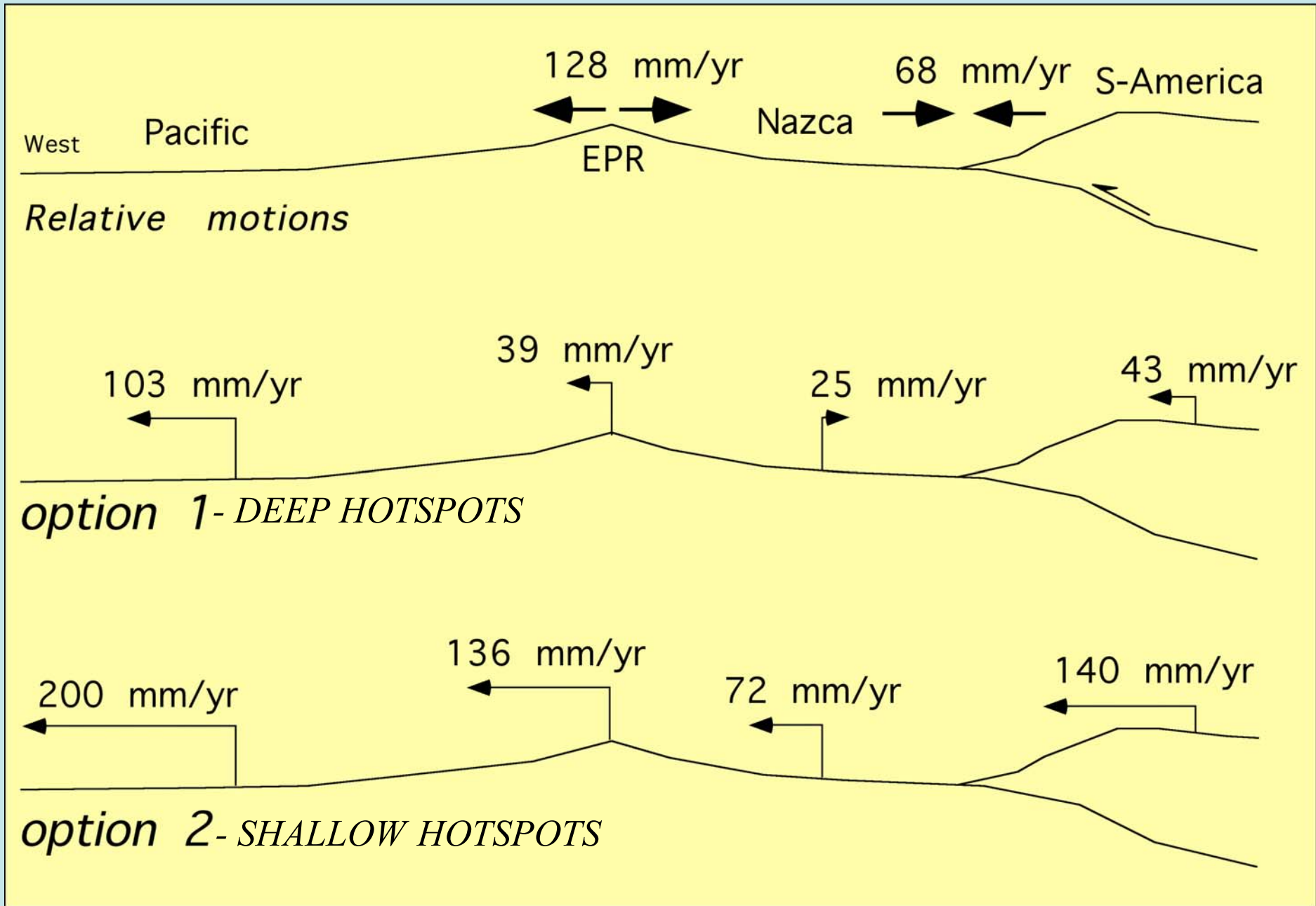
DEEP MANTLE SOURCE

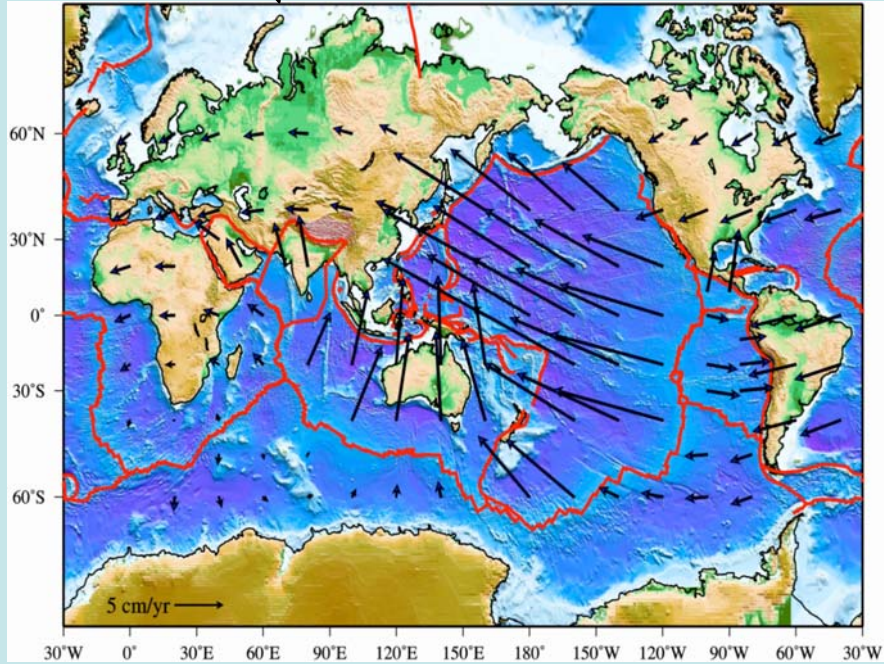
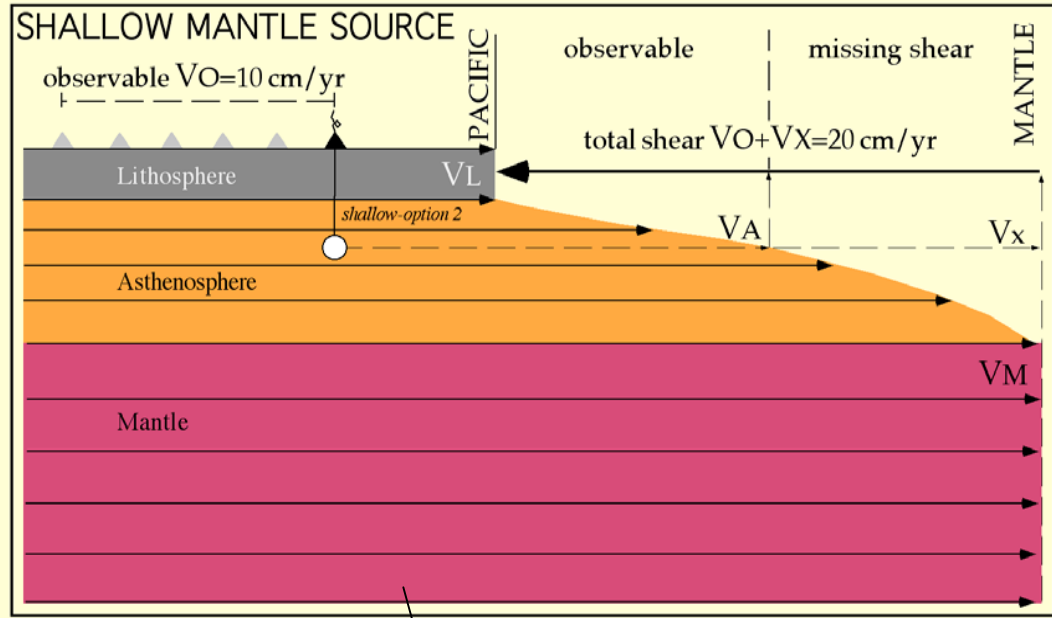
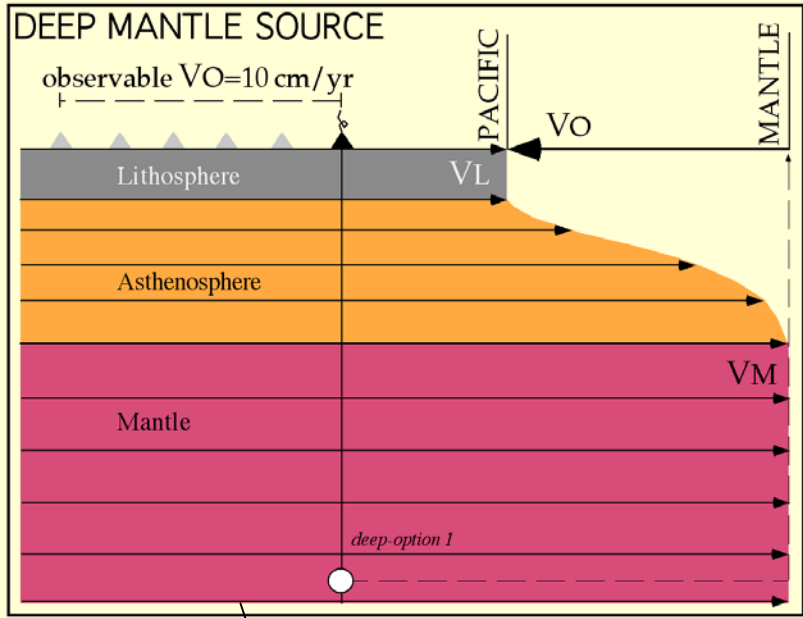
observable $V_O = 10 \text{ cm/yr}$



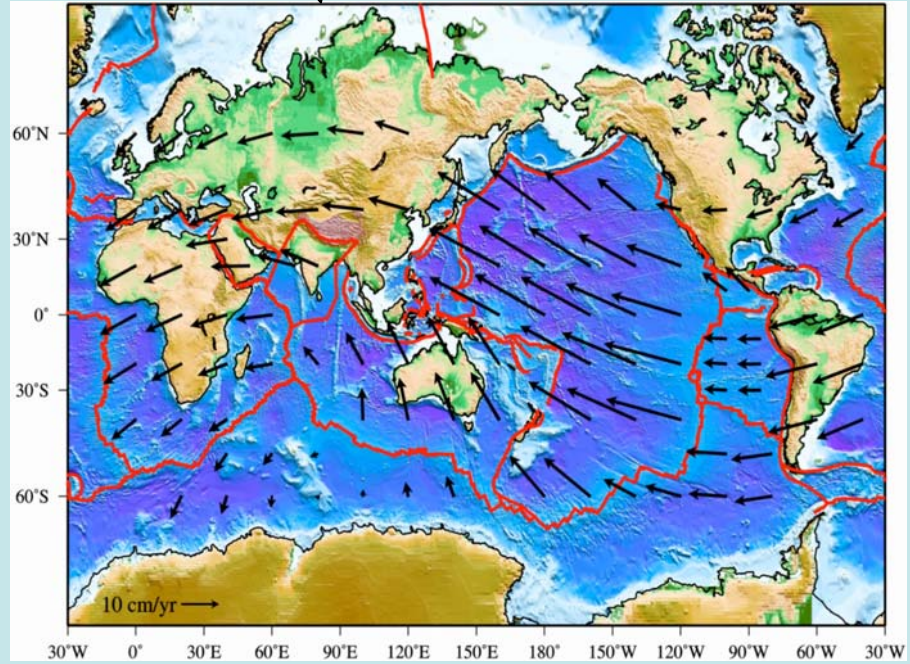
SHALLOW MANTLE SOURCE





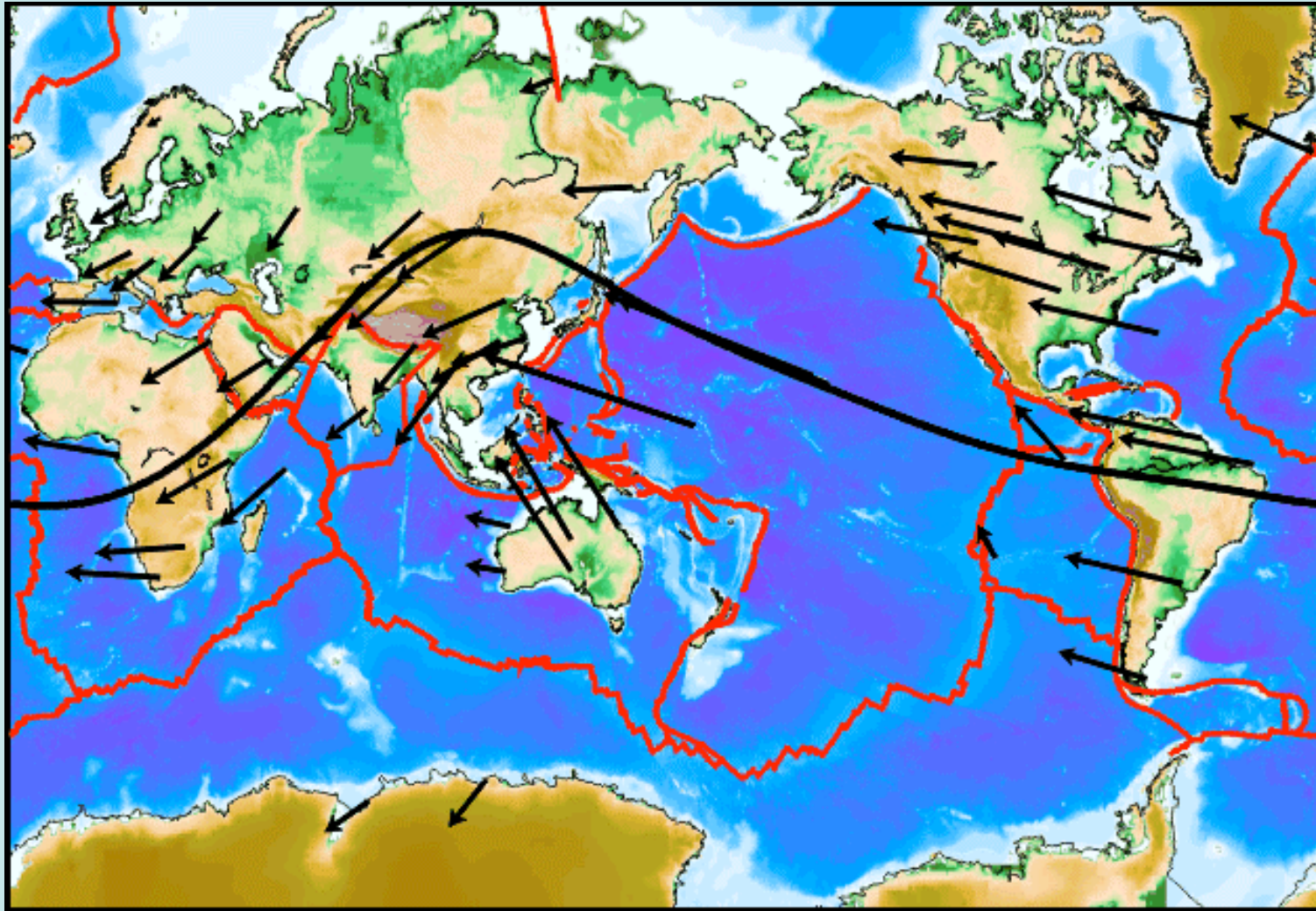


DEEP PLUMES



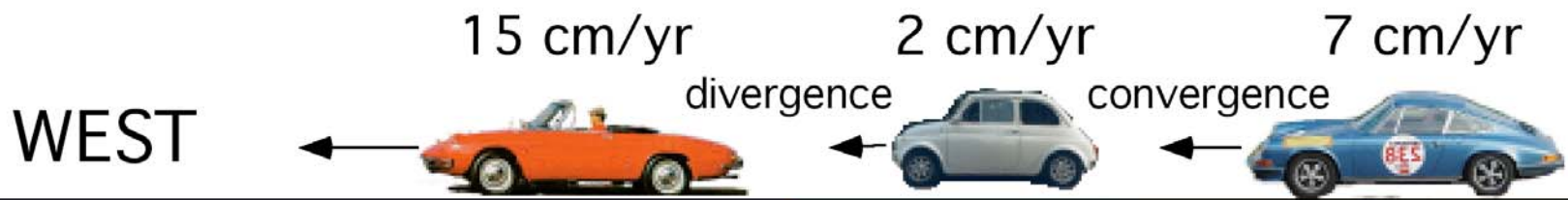
SHALLOW PLUMES

Tectonic “equator”



The sub-lithospheric mantle should move “E-ward”

Cuffaro, 2006



subduction

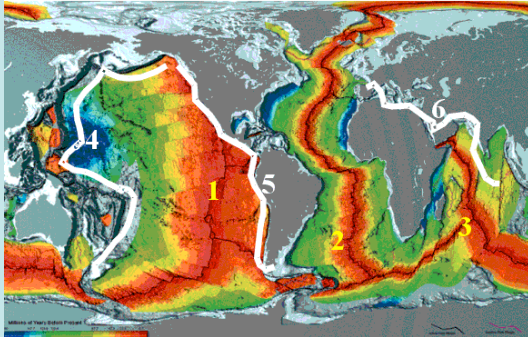
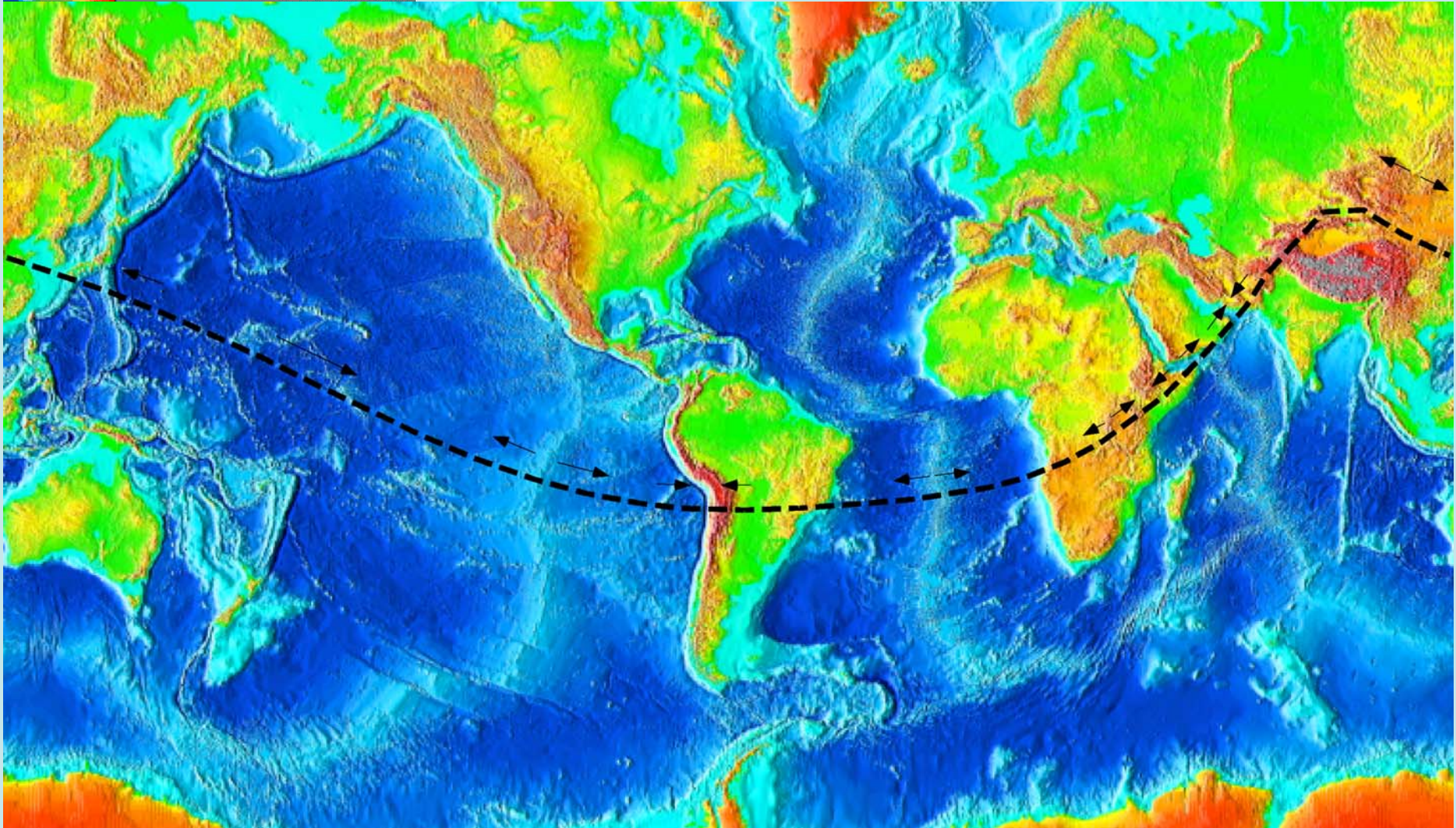


Plate motions are not random during the last 100 Ma, but along a sinusoidal flow:
TECTONIC MAINSTREAM



analytical definition of the *tectonic mainstream*

"The line approximating all the Eulerian equators of the crossed plates"

A 3rd order Fourier series in geographic coordinates (φ, λ), may describe the *tectonic mainstream*

after

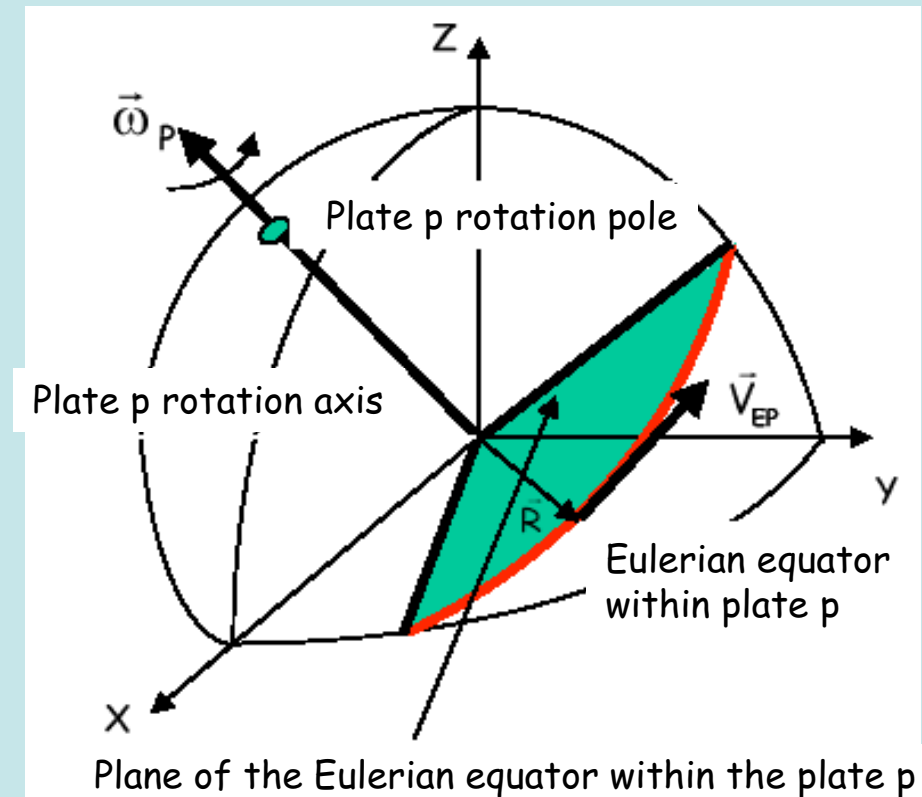
the estimation of the Eulerian equators of crossed plates

$$\varphi = \frac{a_0}{2} + \sum_{i=1}^3 [a_i \cos(i\lambda) + b_i \sin(i\lambda)]$$

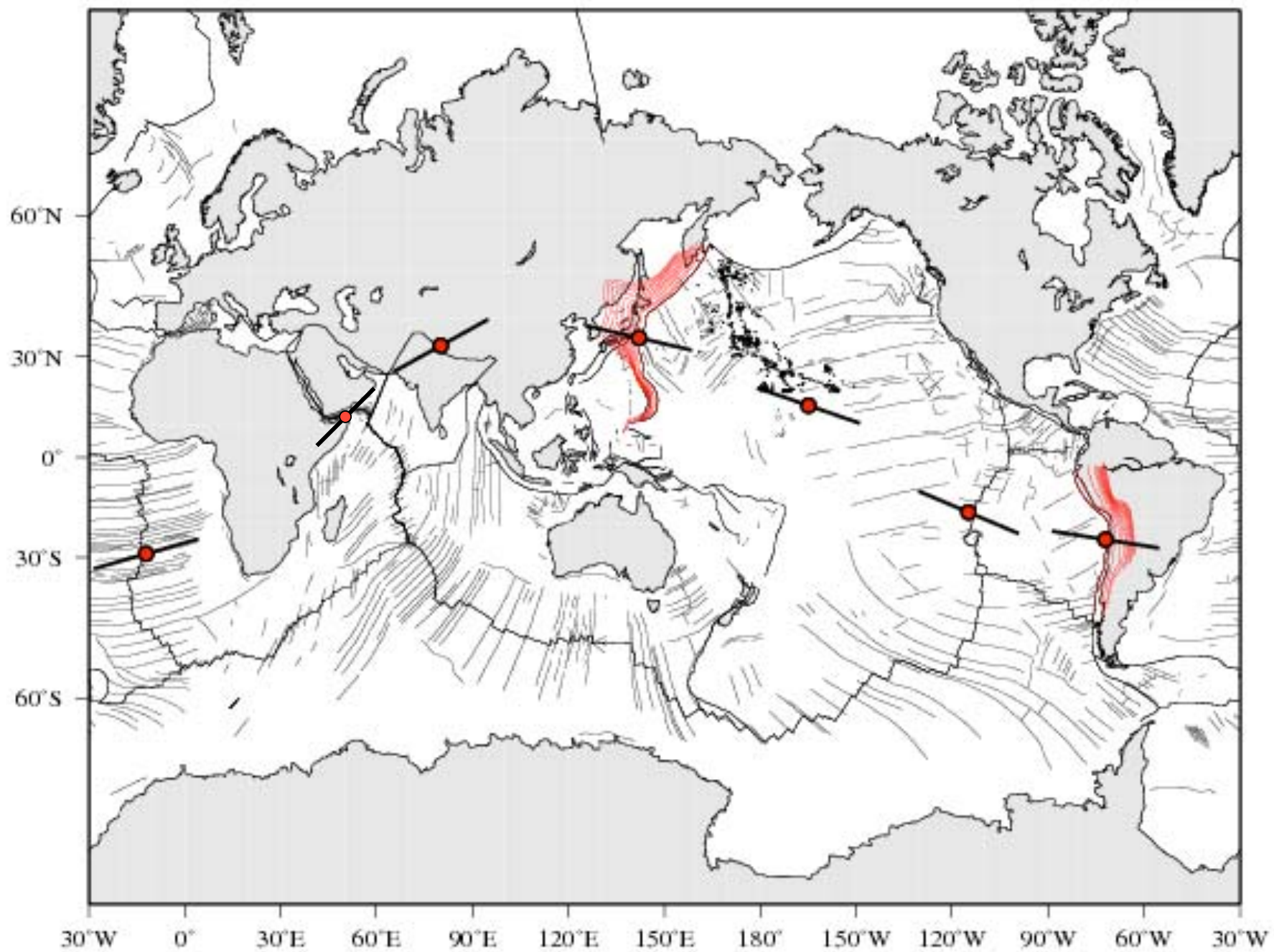
analytical representation of the *tectonic mainstream*

Modeling the “absolute” plate motions assumptions

- Spherical approximation
- “Absolute” plate motions modeled as 3D rotations (deformations neglected)
- Eulerian equator “image” of the *tectonic mainstream* within each crossed plate



$$\omega_{Xp} \cos \varphi \cos \lambda + \omega_{Yp} \cos \varphi \sin \lambda + \omega_{Zp} \sin \varphi = 0$$



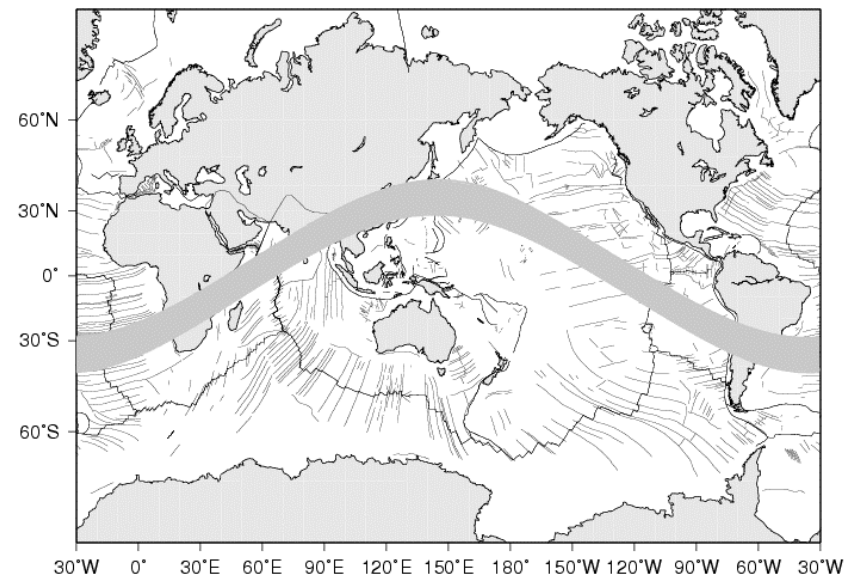
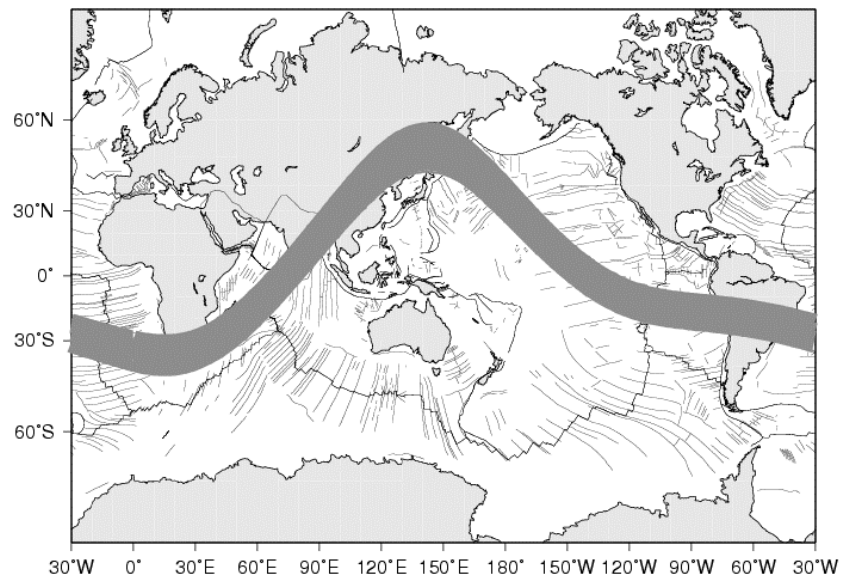
Tectonic mainstream representations latitude confidence intervals (1 sigma)

deep mantle

$\pm 7.3^\circ$

mid asthenosphere

$\pm 6.8^\circ$



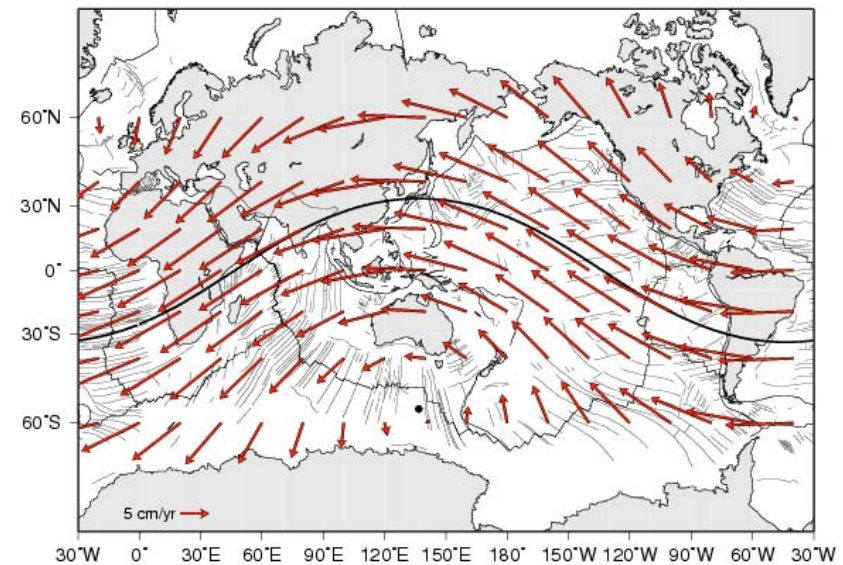
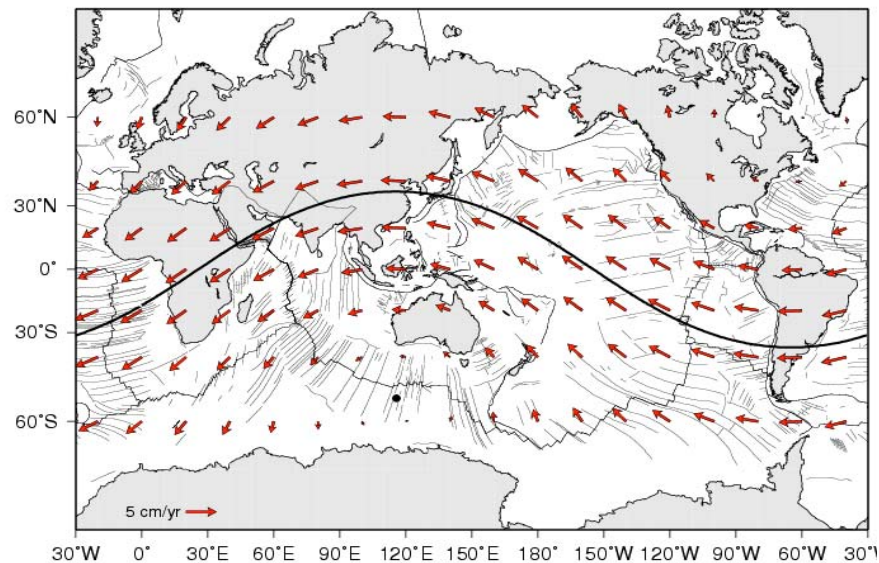
Global lithospheric net rotation tangential velocities and equators

deep mantle

mid asthenosphere

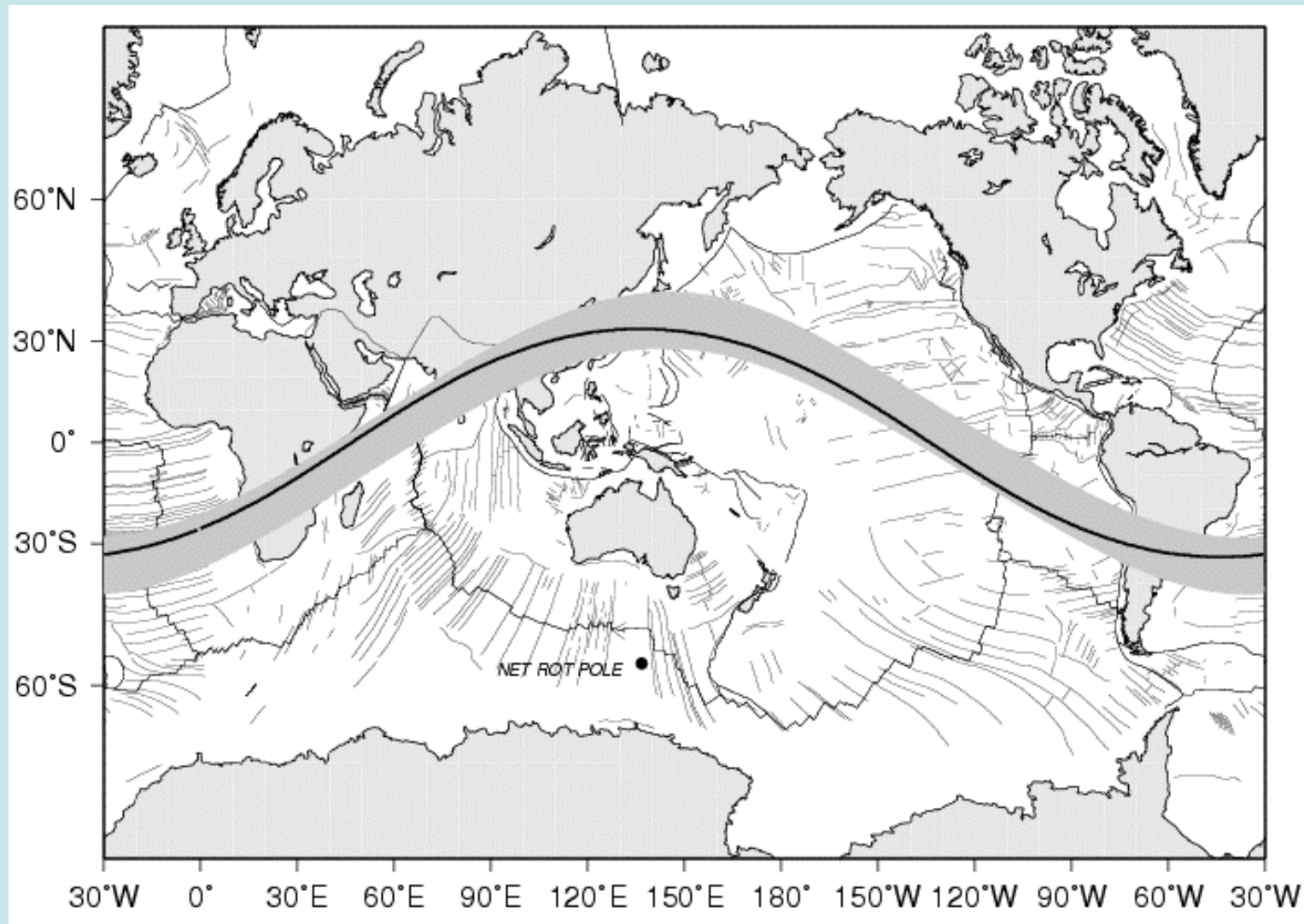
max velocity 4.1 cm/yr

max velocity 13.4 cm/yr

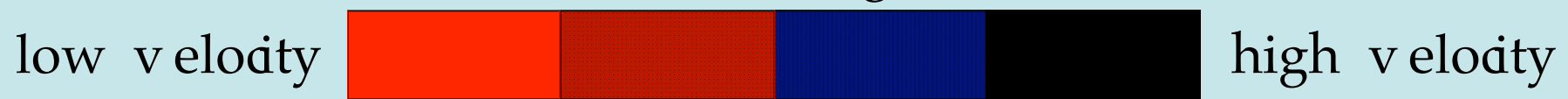
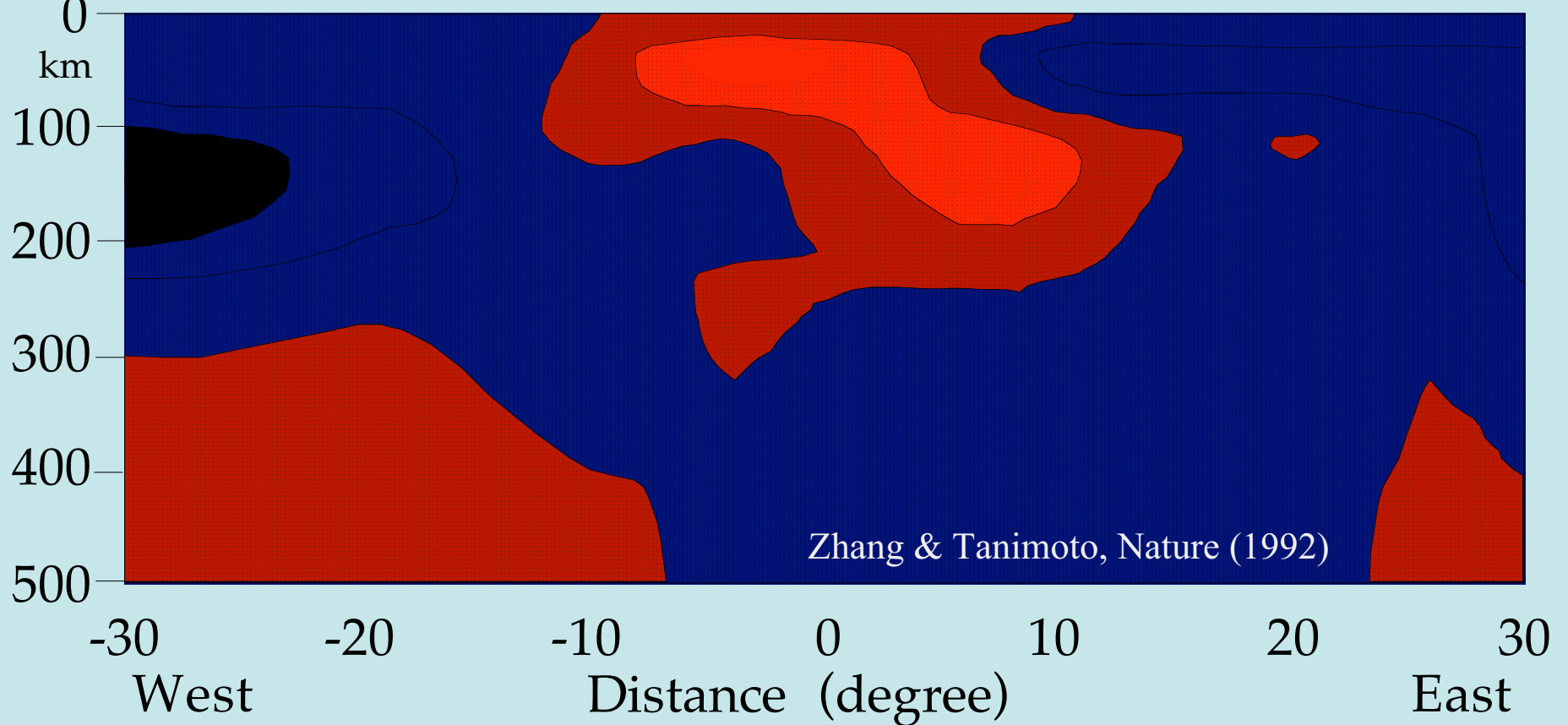
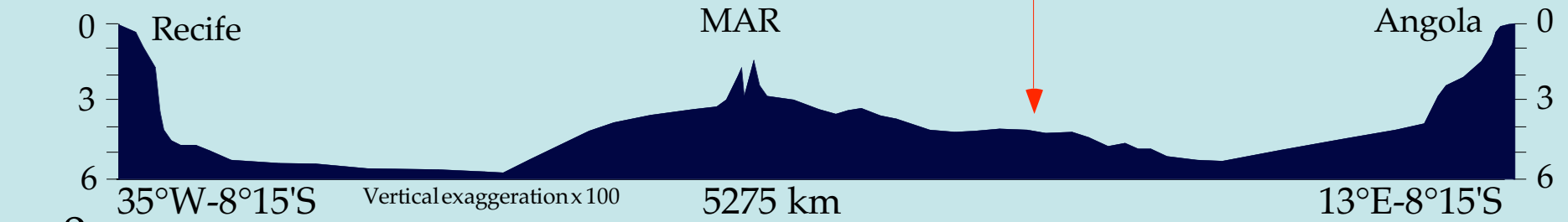


Global lithospheric flow

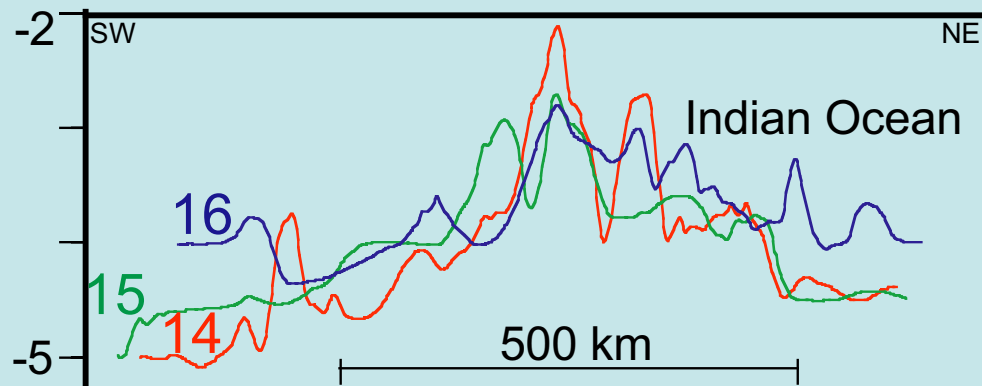
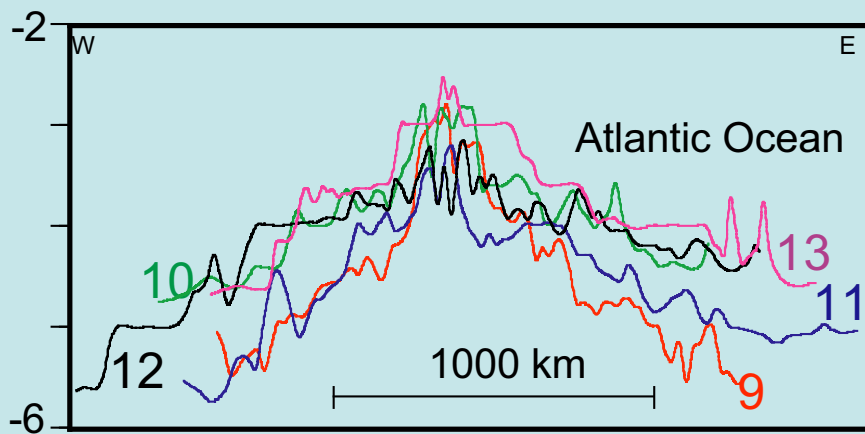
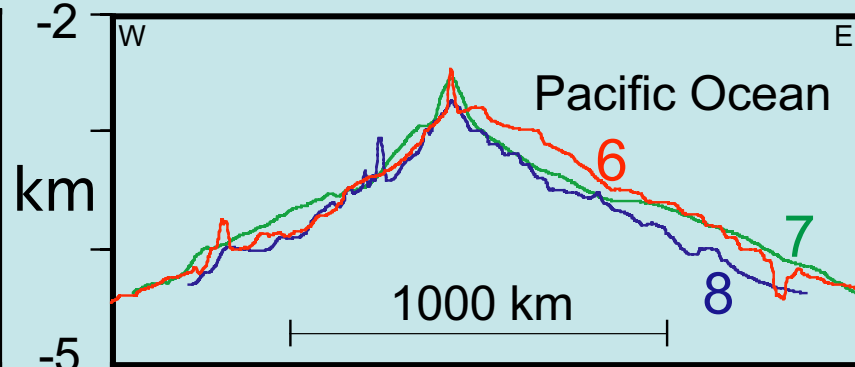
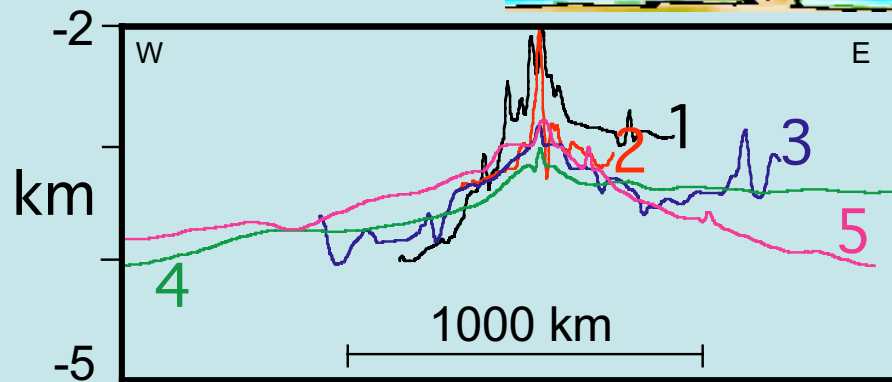
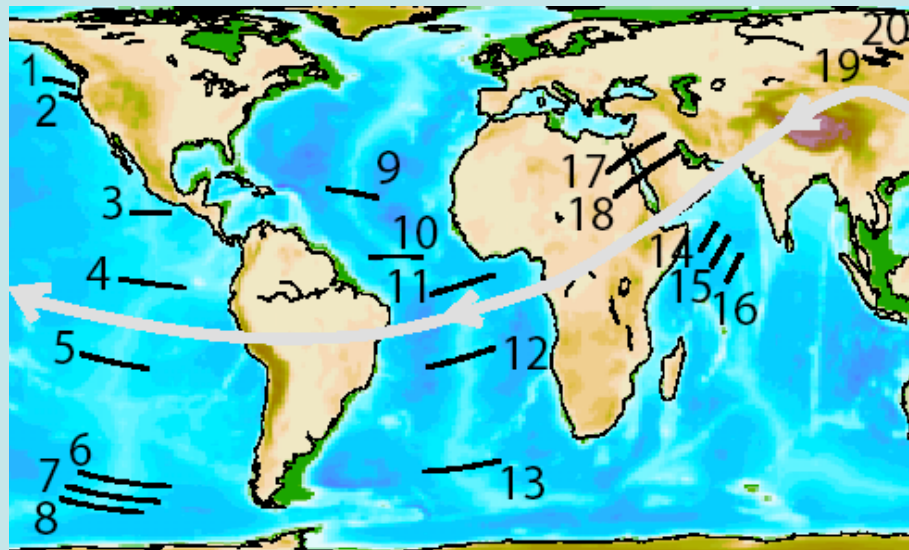
mid astenosphere net rotation equator lies within the *tectonic mainstream* latitude band



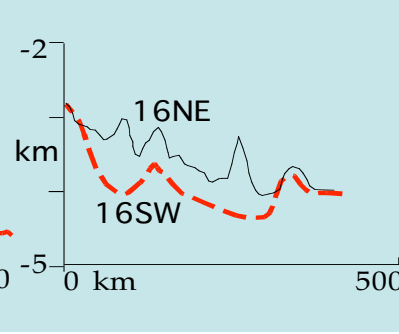
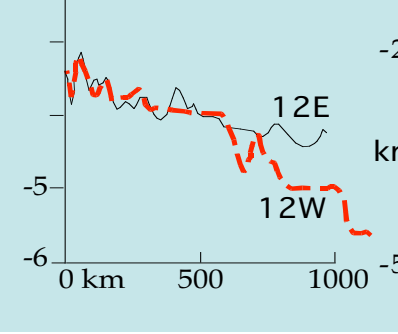
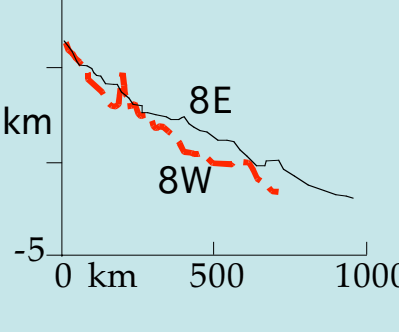
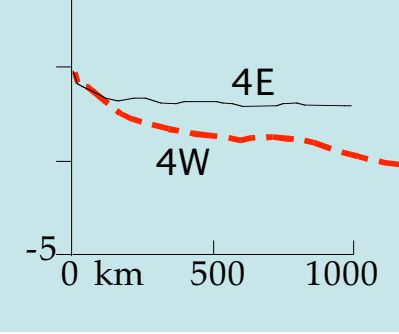
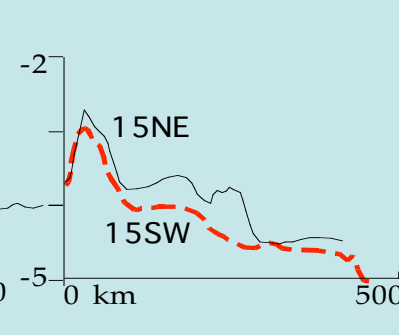
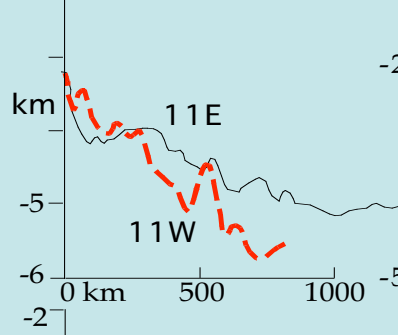
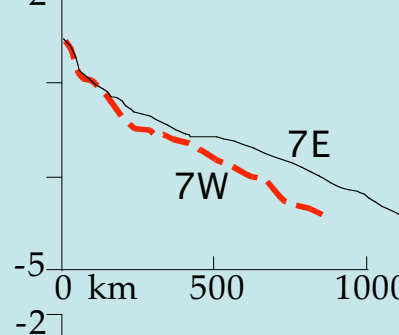
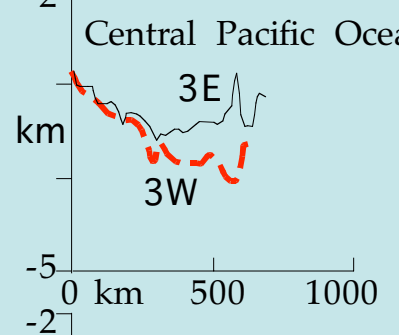
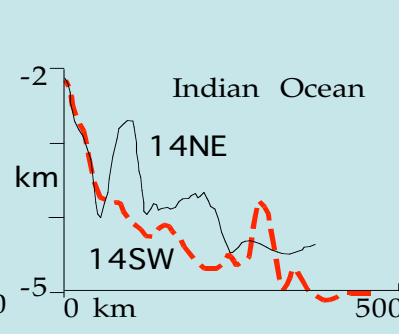
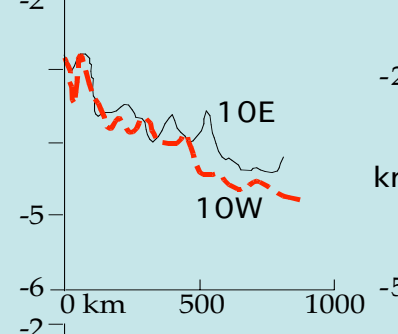
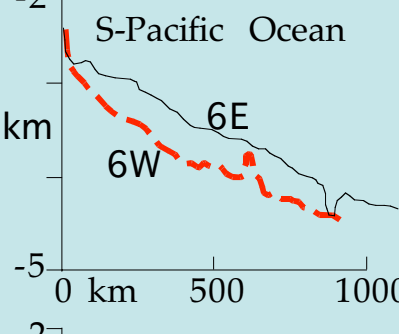
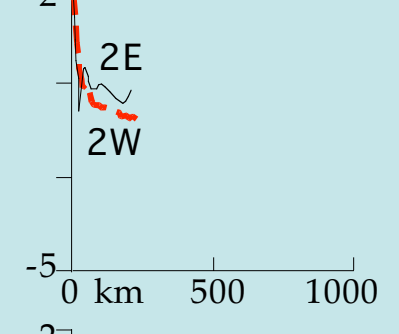
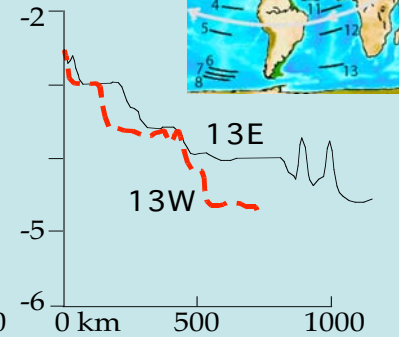
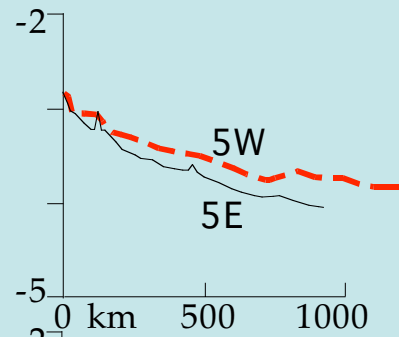
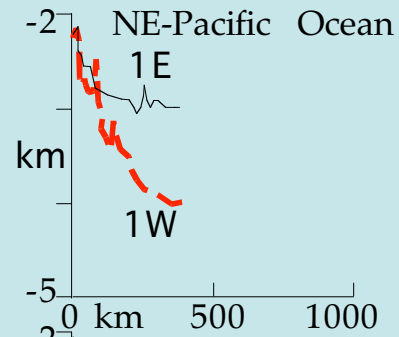
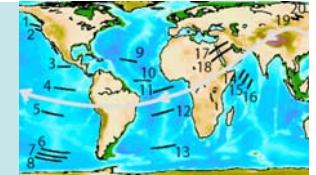
Shallower and less steep flank



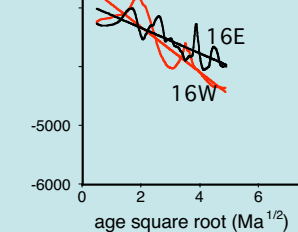
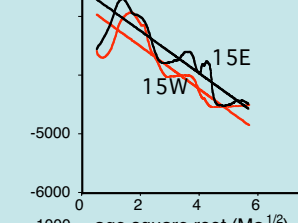
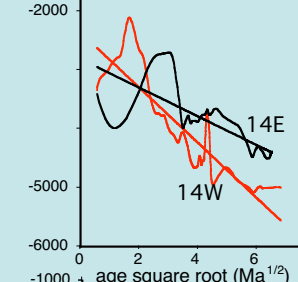
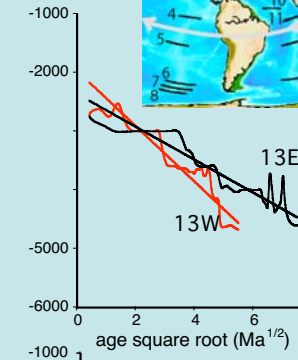
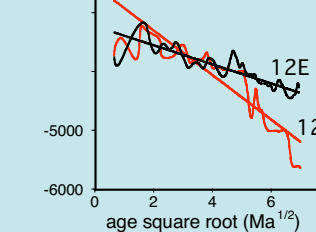
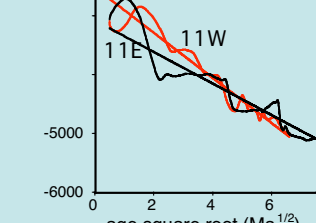
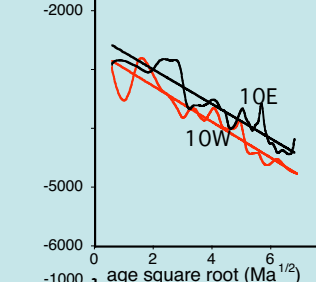
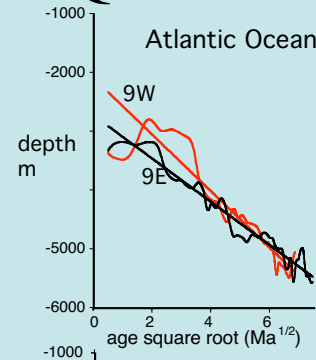
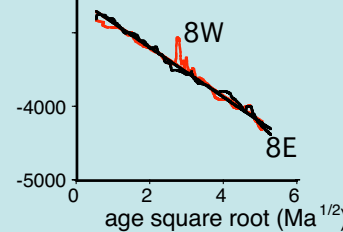
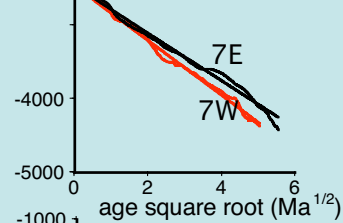
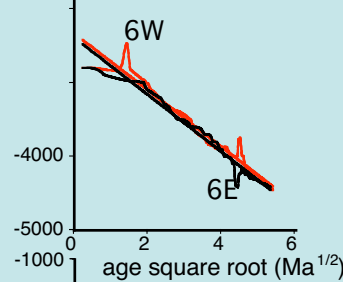
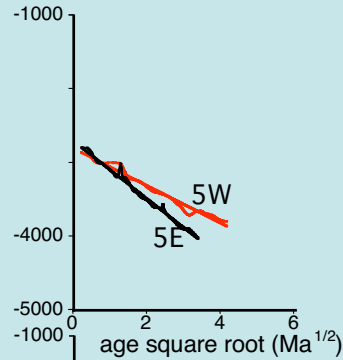
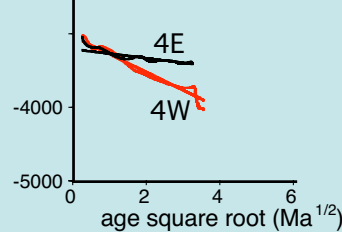
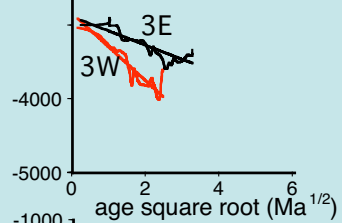
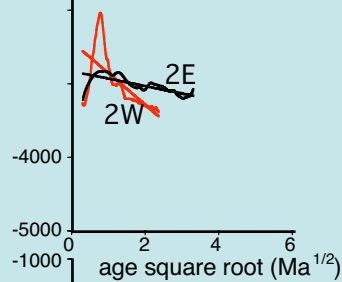
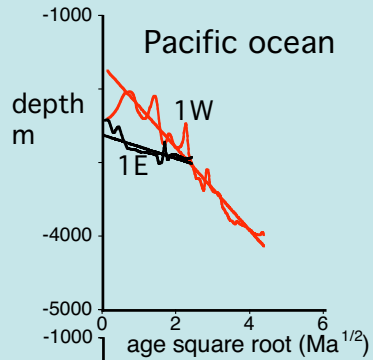
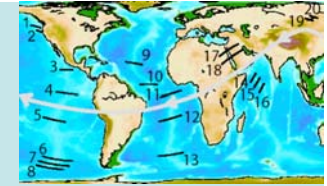
S. MidAtlantic Ridge (-21 < Lat < -11)

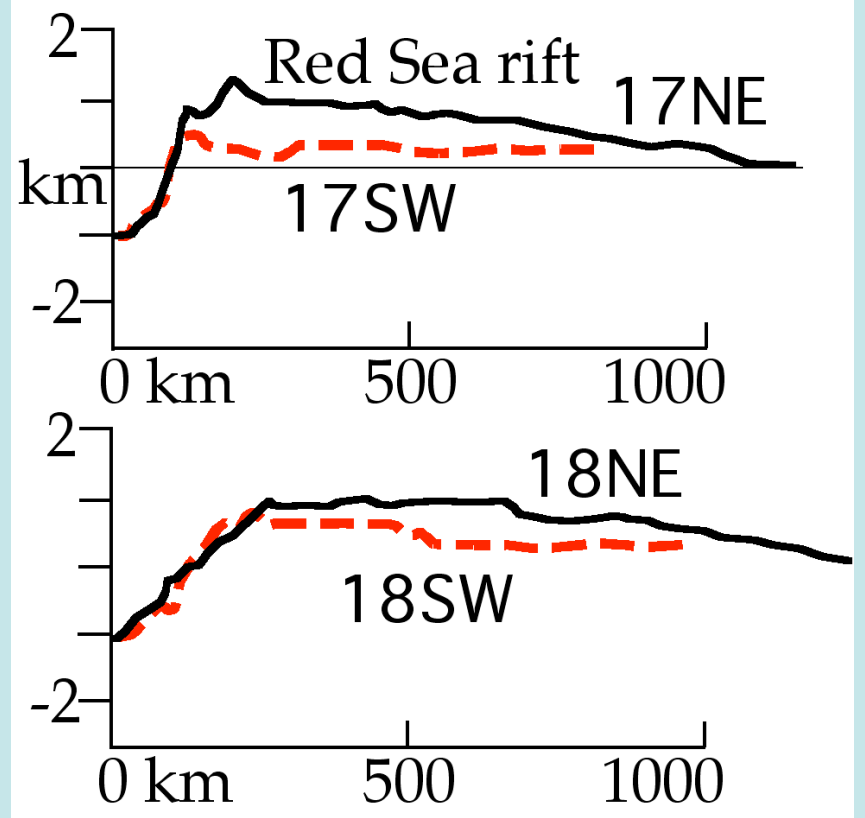
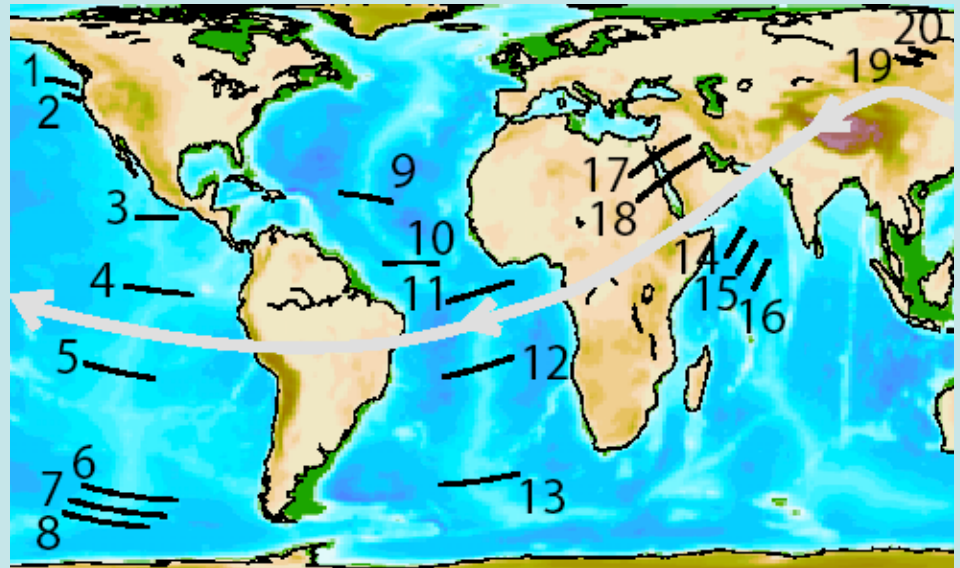
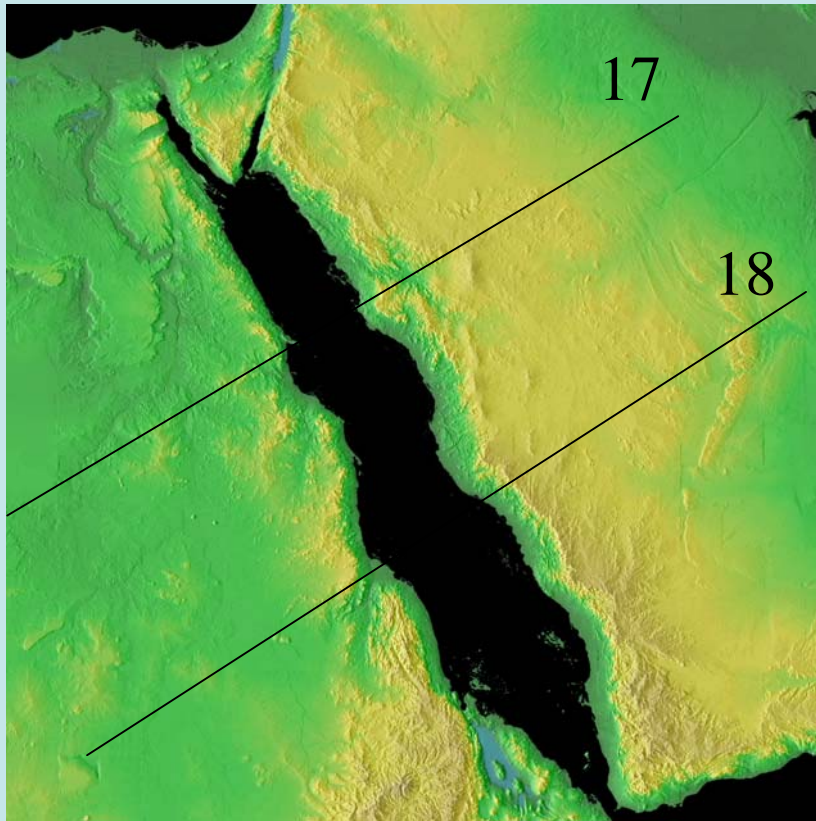


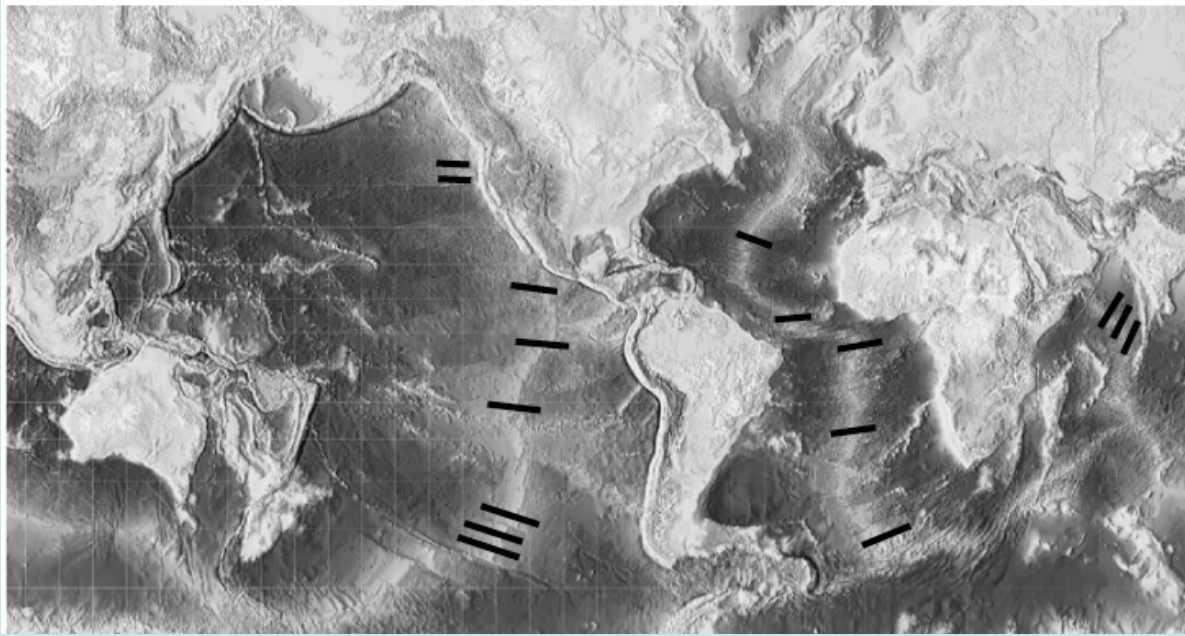
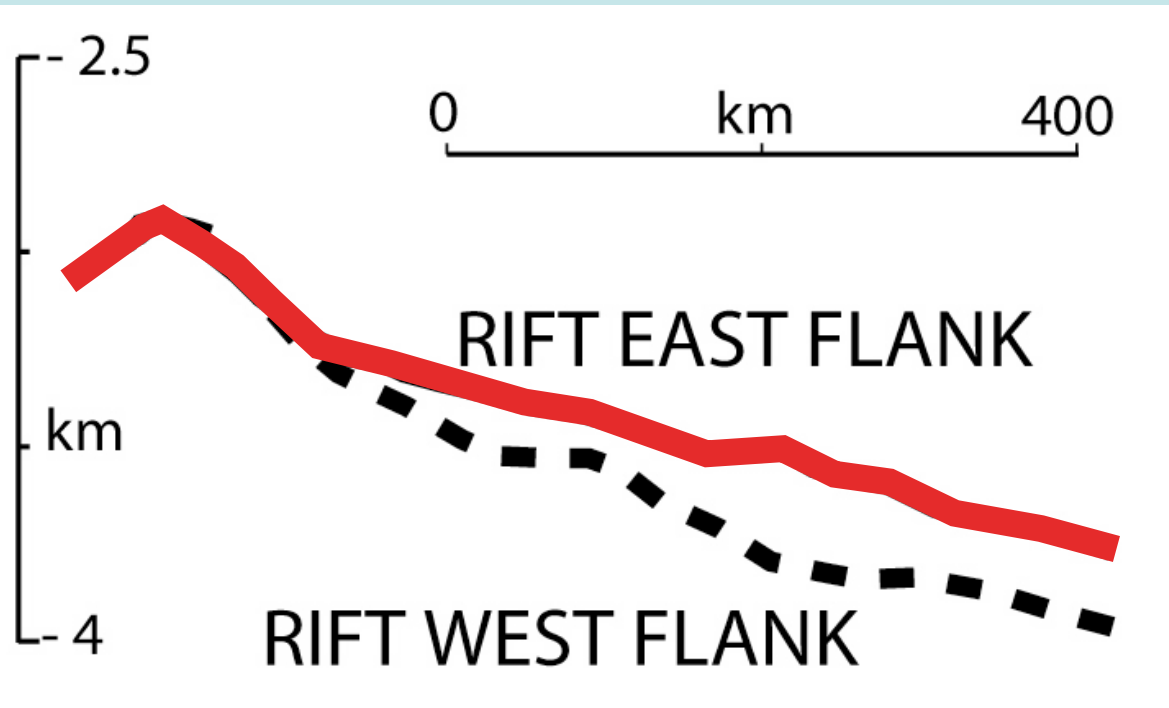
BATHYMETRY vs DISTANCE



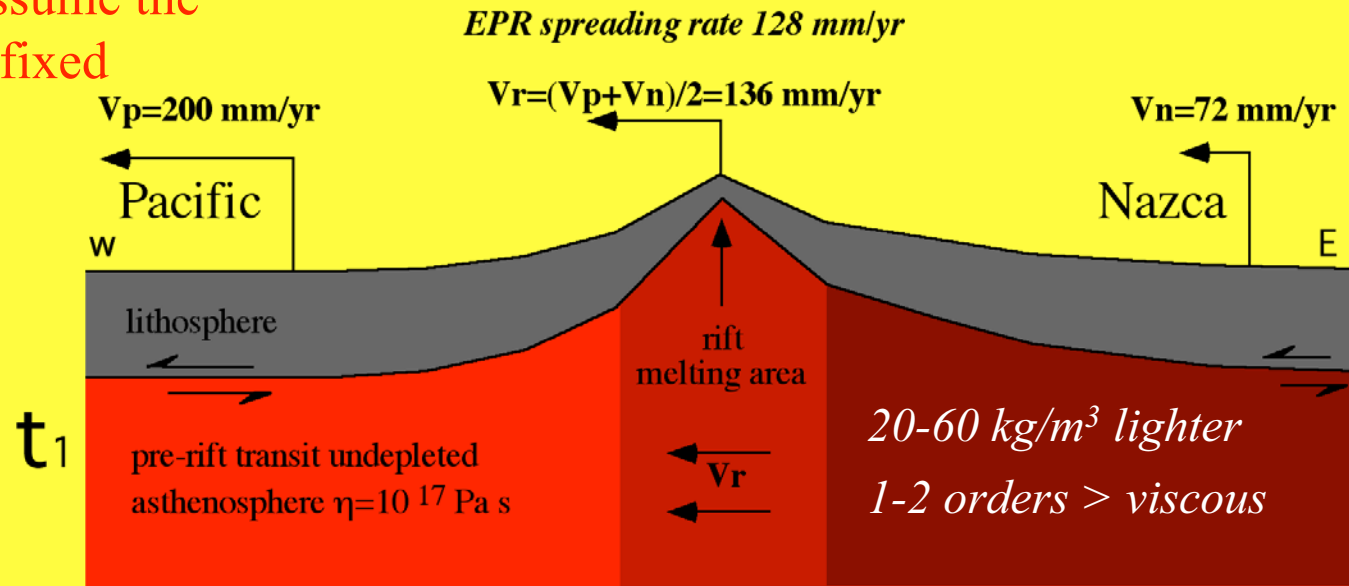
BATHYMETRY vs AGE SQUARE ROOT







Let's assume the mantle fixed

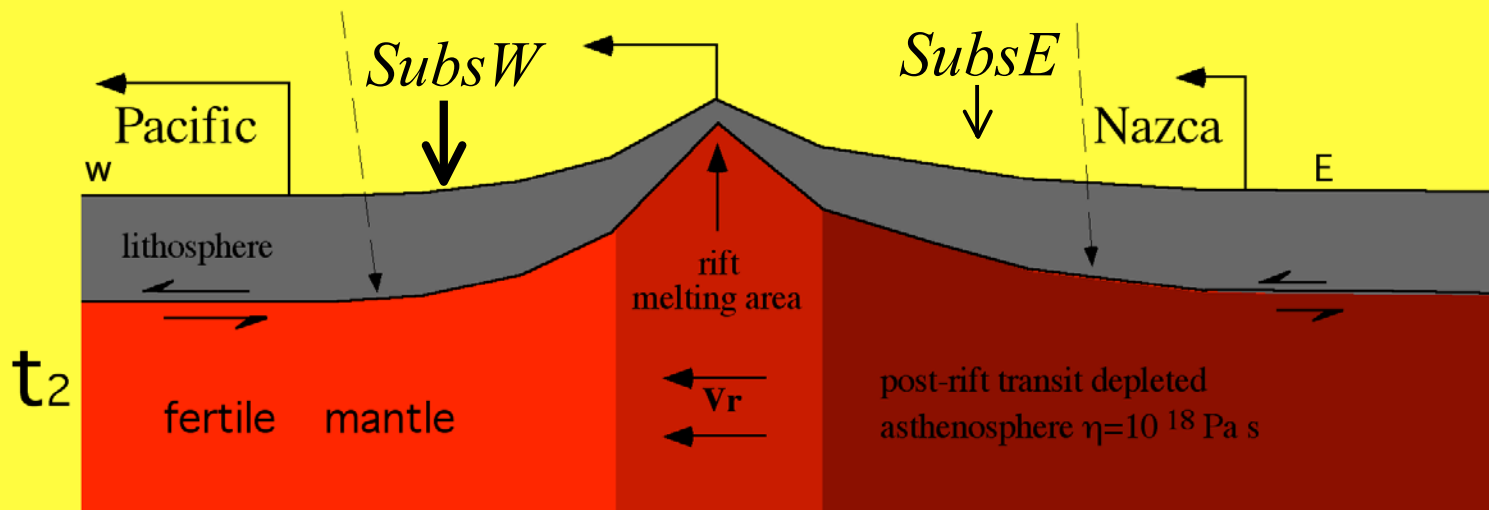


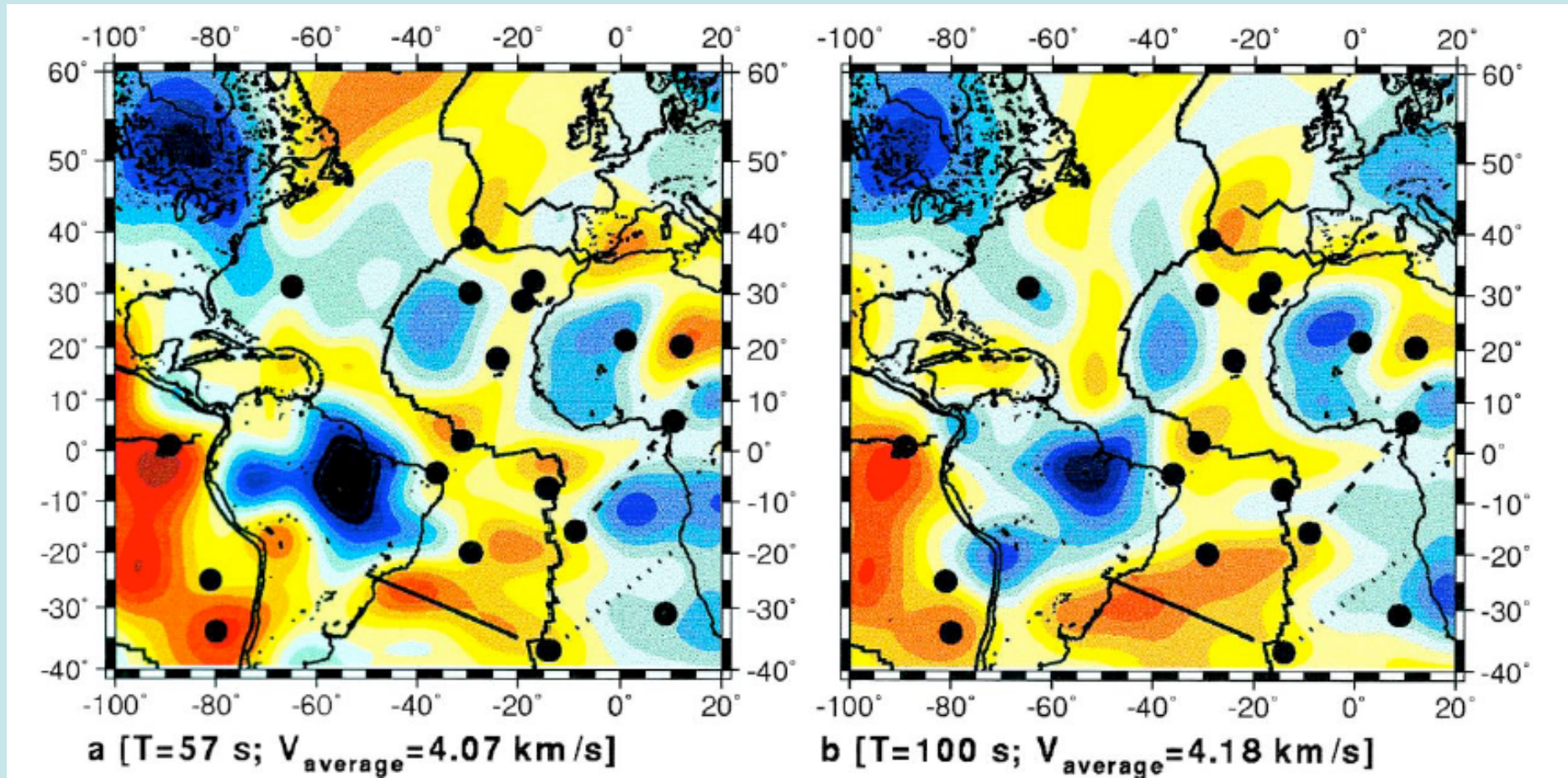
$SubsW > SubsE$

Melting area moving west at V_r

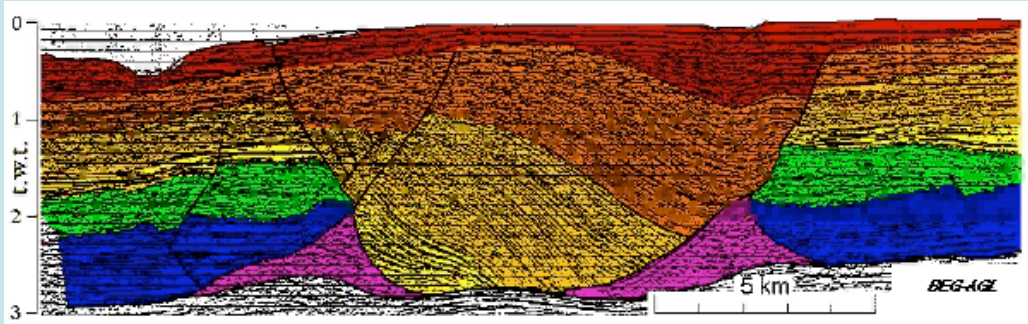
FASTER DECOLLEMENT IN THE WESTERN LIMB

higher viscosity = more coupled contact (eastern plate steadily slower)

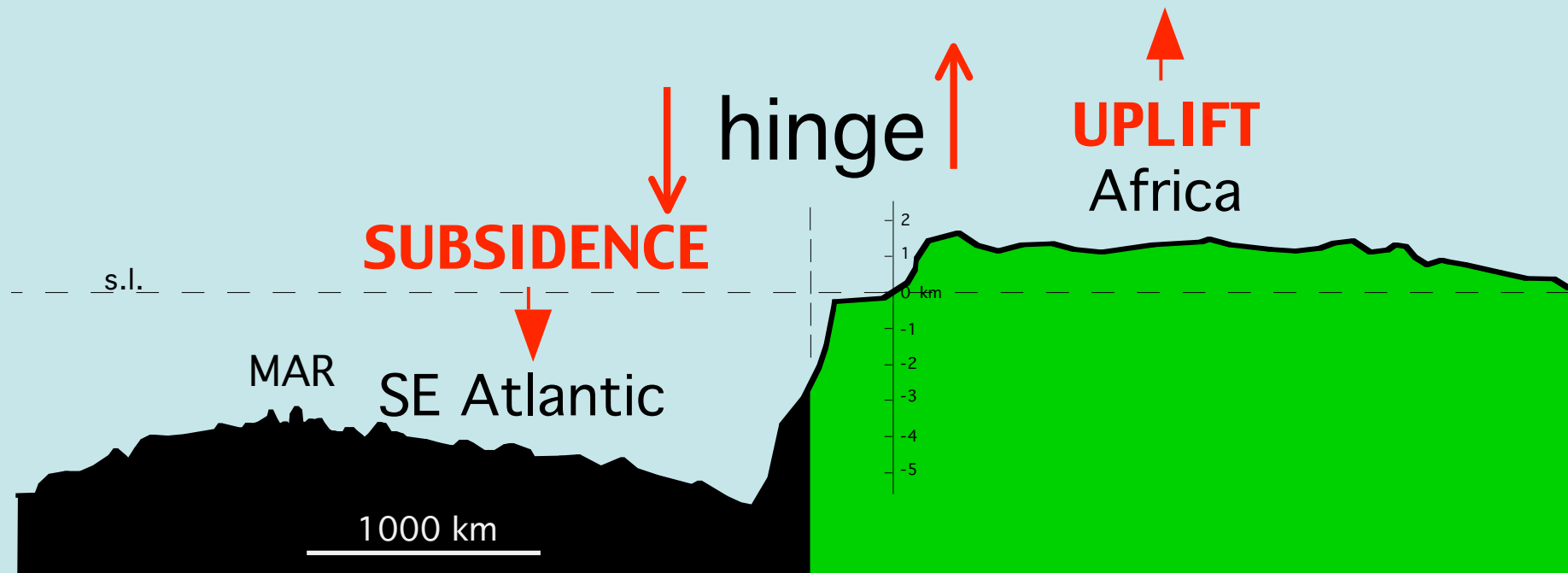




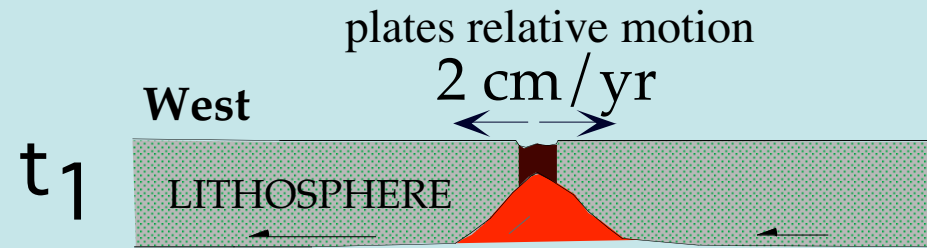
Silveira et al., 1998, PEPI



- Different isostatic answer due to lateral variation in composition and temperature of the lithosphere
- Coherent system: no independent mantle plume confined underneath the continent



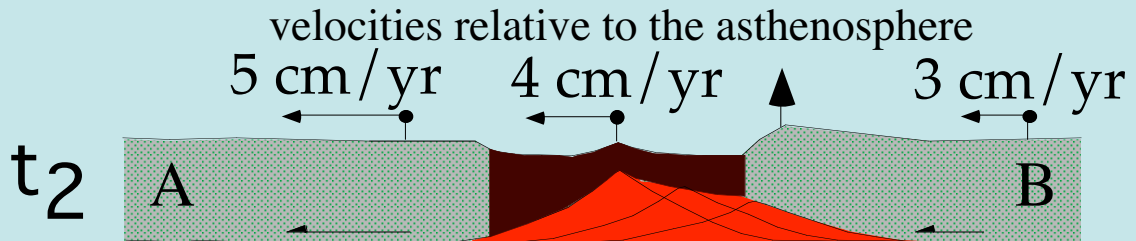
rifting model



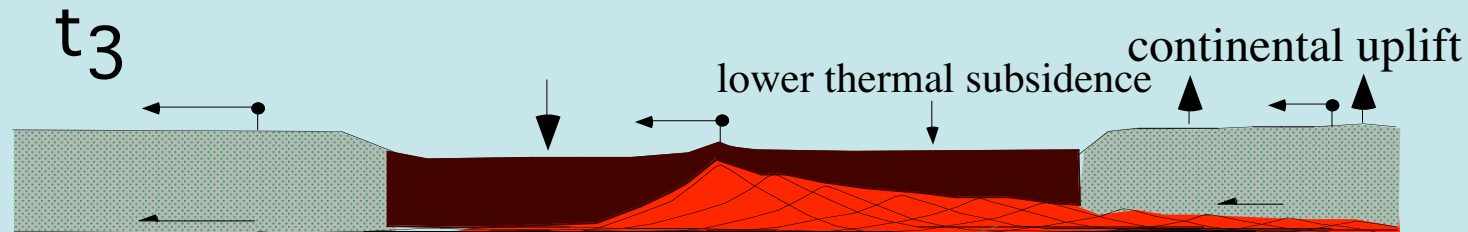
lower density anomaly



$Ridge\ speed = (V_A + V_B) / 2$

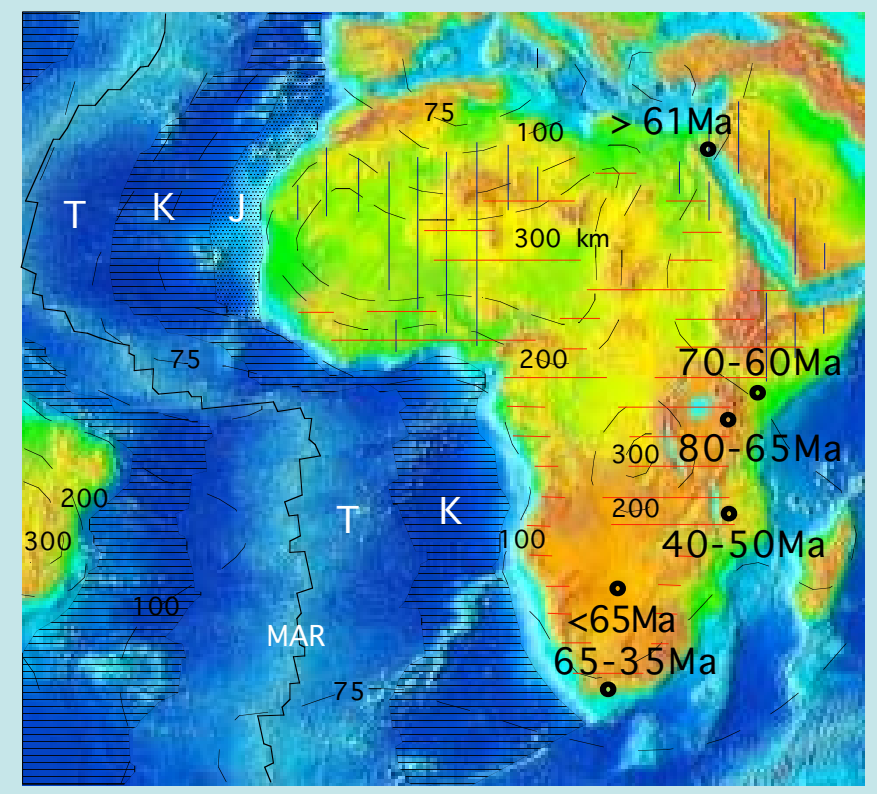
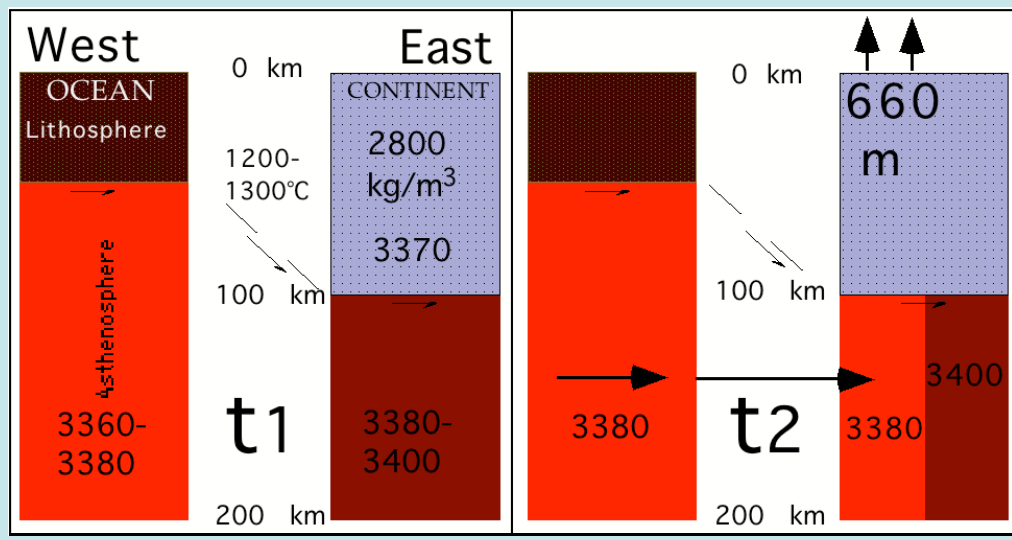
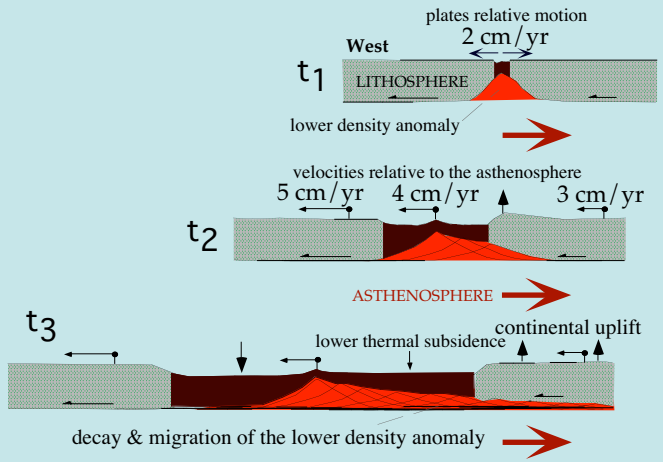


ASTHENOSPHERE



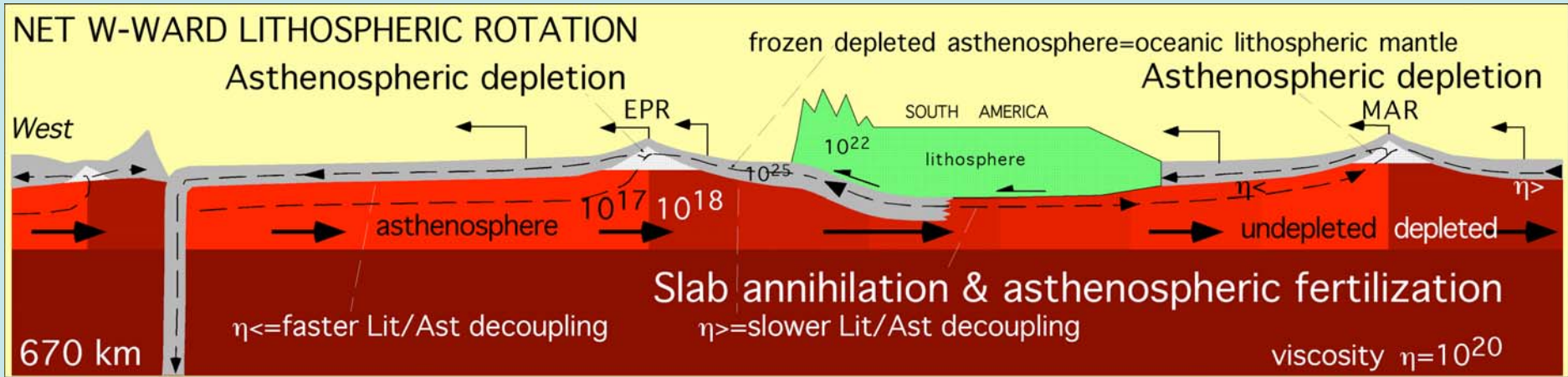
decay & migration of the lower density anomaly





...and Europe?

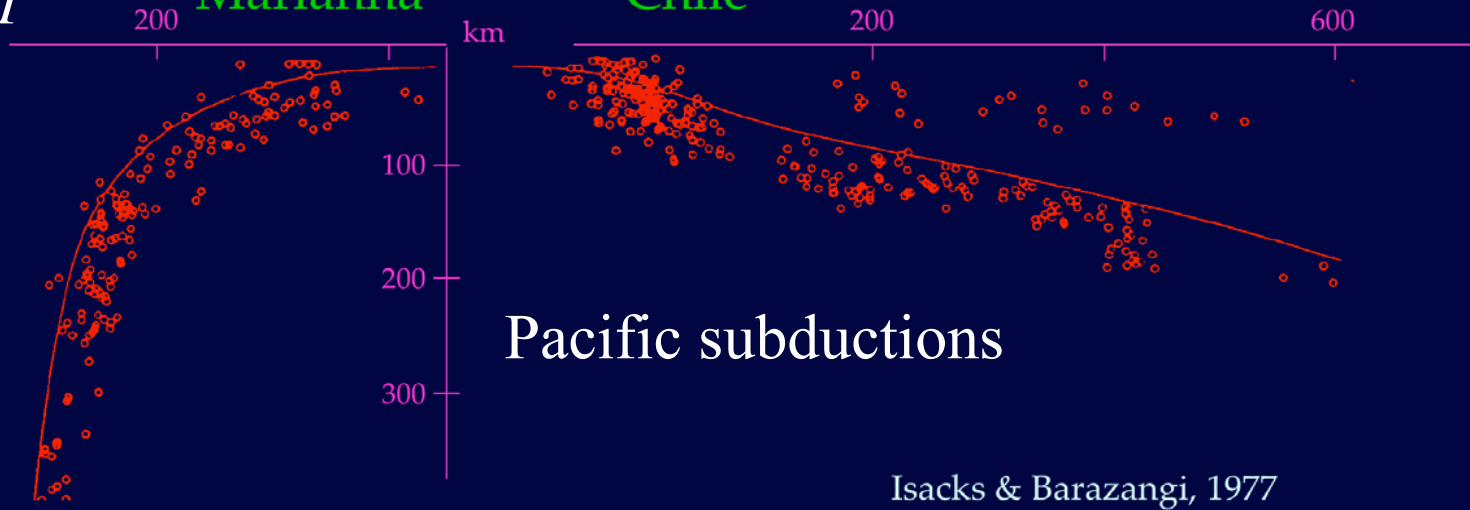




WEST

Marianna

Chile

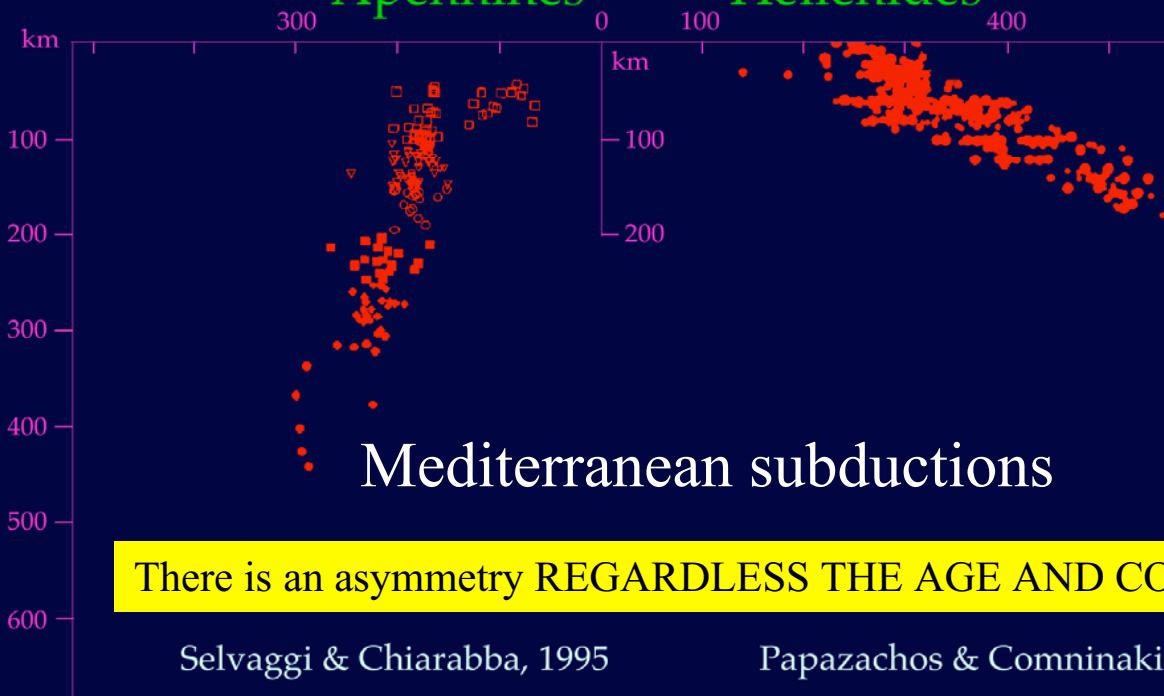


Pacific subductions

Isacks & Barazangi, 1977

Apennines

Hellenides

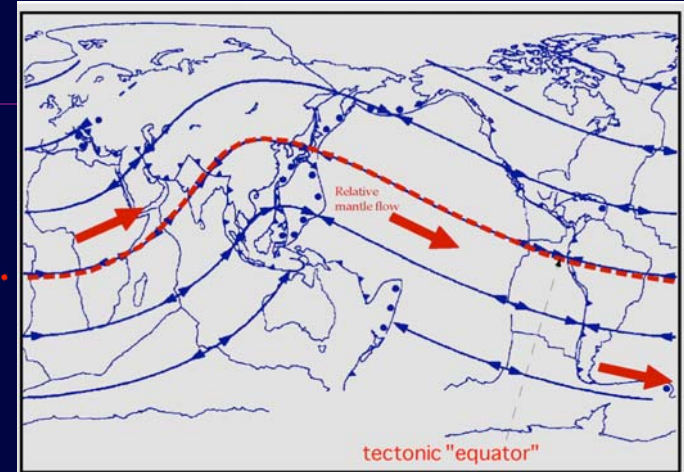


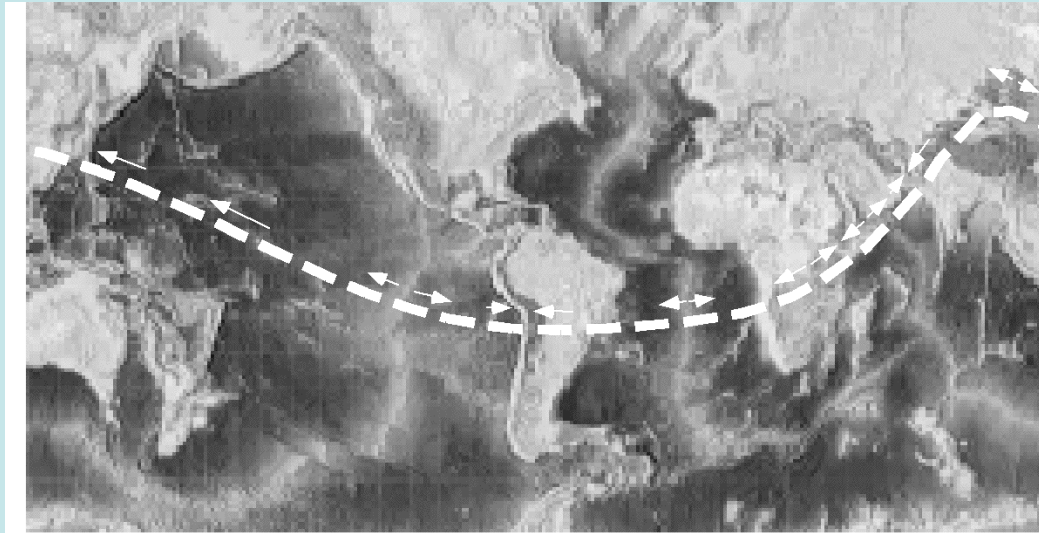
Mediterranean subductions

There is an asymmetry REGARDLESS THE AGE AND COMPOSITION of the lithosphere

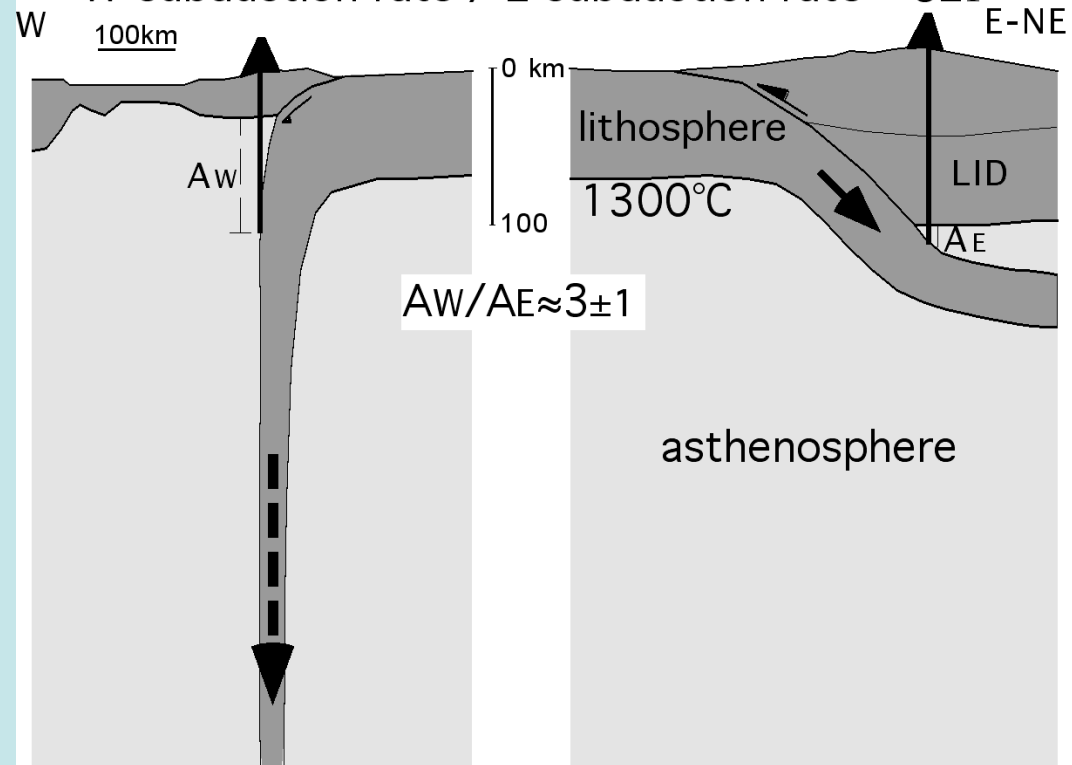
Selvaggi & Chiarabba, 1995

Papazachos & Comninakis, 1977





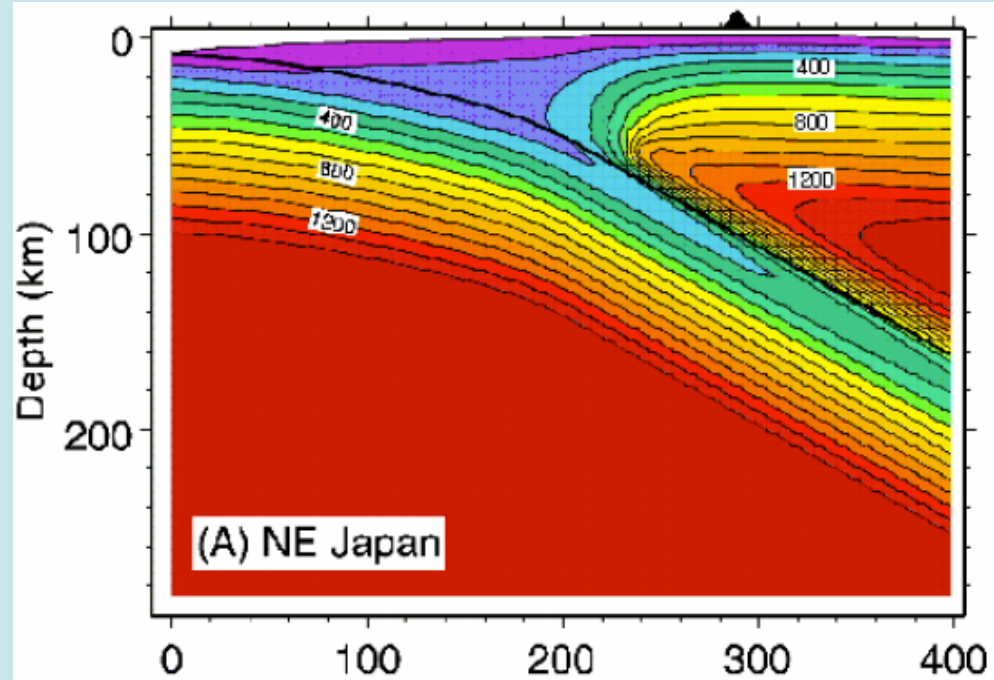
W-subduction rate / E-subduction rate $\approx 3 \pm 1$



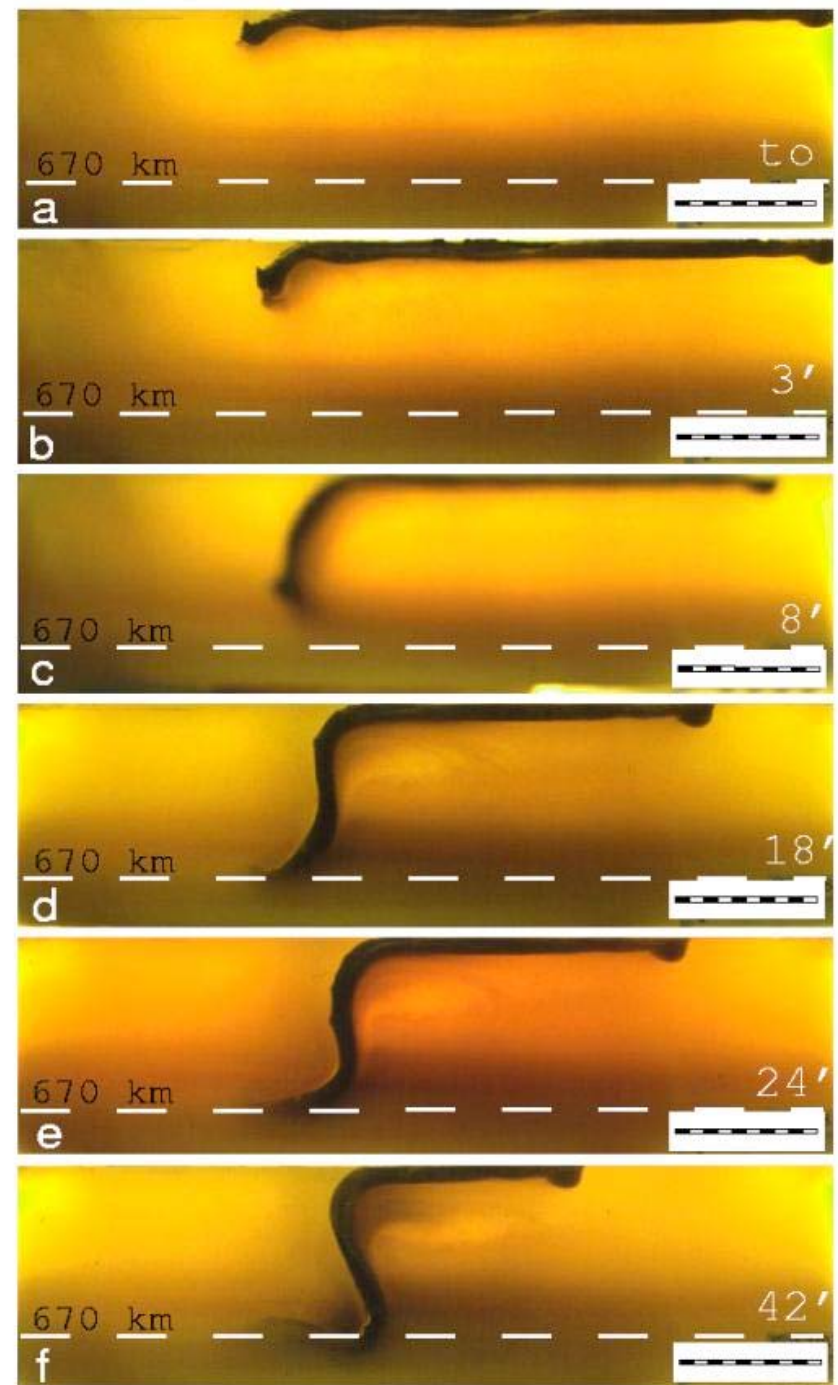


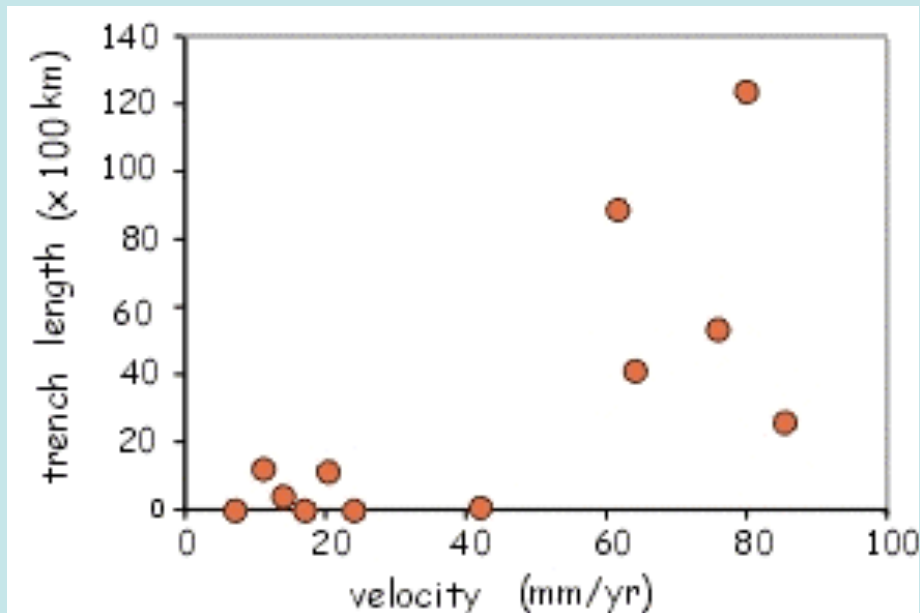
WHERE IS THE SLAB PULL?

Is the slab a priori heavier, or it becomes denser due to phase transitions induced by the drag of the mantle?



$$\text{Slab pull} = 10^{13} \text{ N m}^{-1}$$





NNR

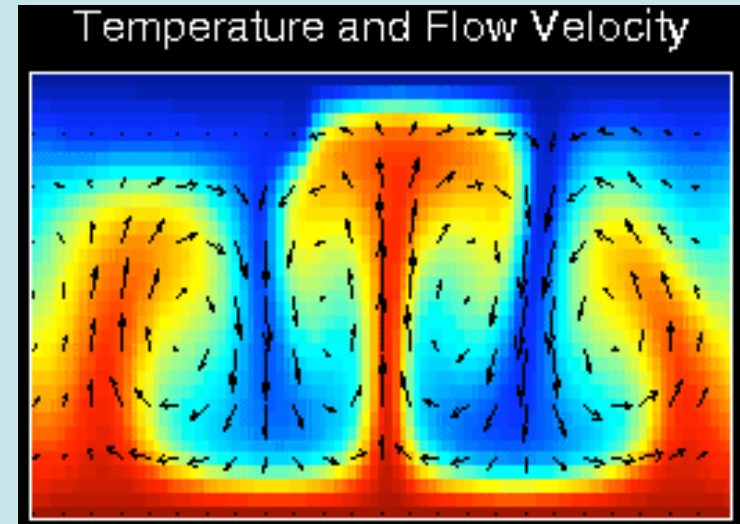
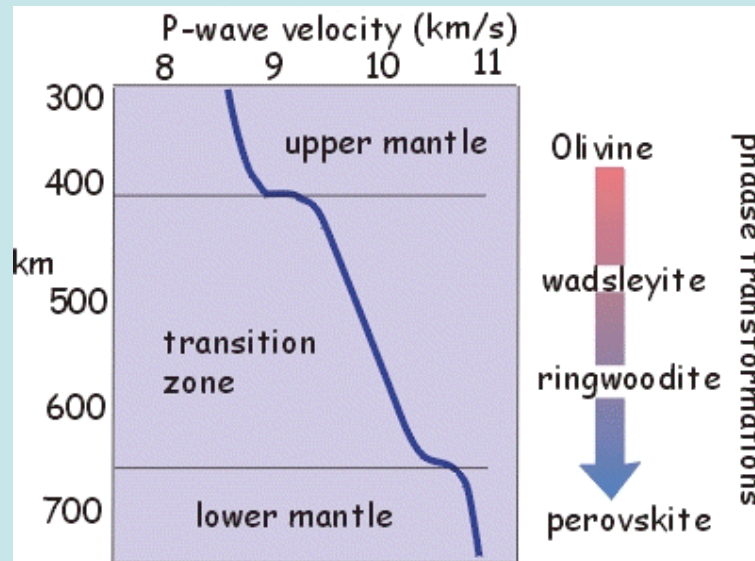


HSRF



The observations that plates with large slabs are faster (e.g. Forsyth & Uyeda, 1975) is a circular reasoning:

- 1) they could have longer subduction zones because they are faster, not necessarily vice-versa;
- 2) but are they faster relative to what?
- 3) Nazca has fast convergence rates in a not net rotation frame, which is an artifact, but it is much slower relative to the mantle



Assumptions:
Comments

The Mantle is compositionally homogeneous:

*Very unlikely, all Earth is intensely stratified by density from the topmost atmosphere down to the core;
If the mantle was homogeneous and movements are only driven by thermal contrast, why the LID does not detach in the middle of a plate?*

Uprising mantle is laterally accompanied by down-welling:

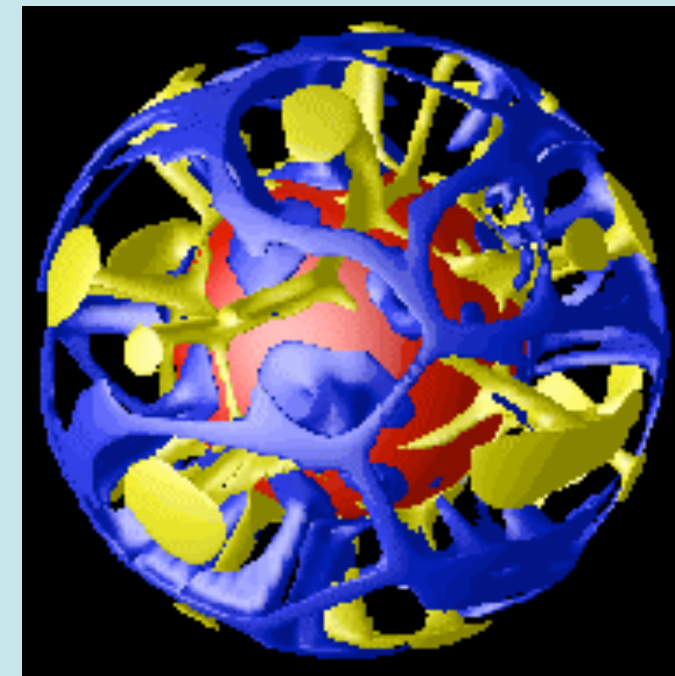
Atlantic, E-Africa and Indian rifts have not intervening subductions; there are also several cases of paired subduction zones without rifts in between

Uprising and down-welling mantle currents are stationary:

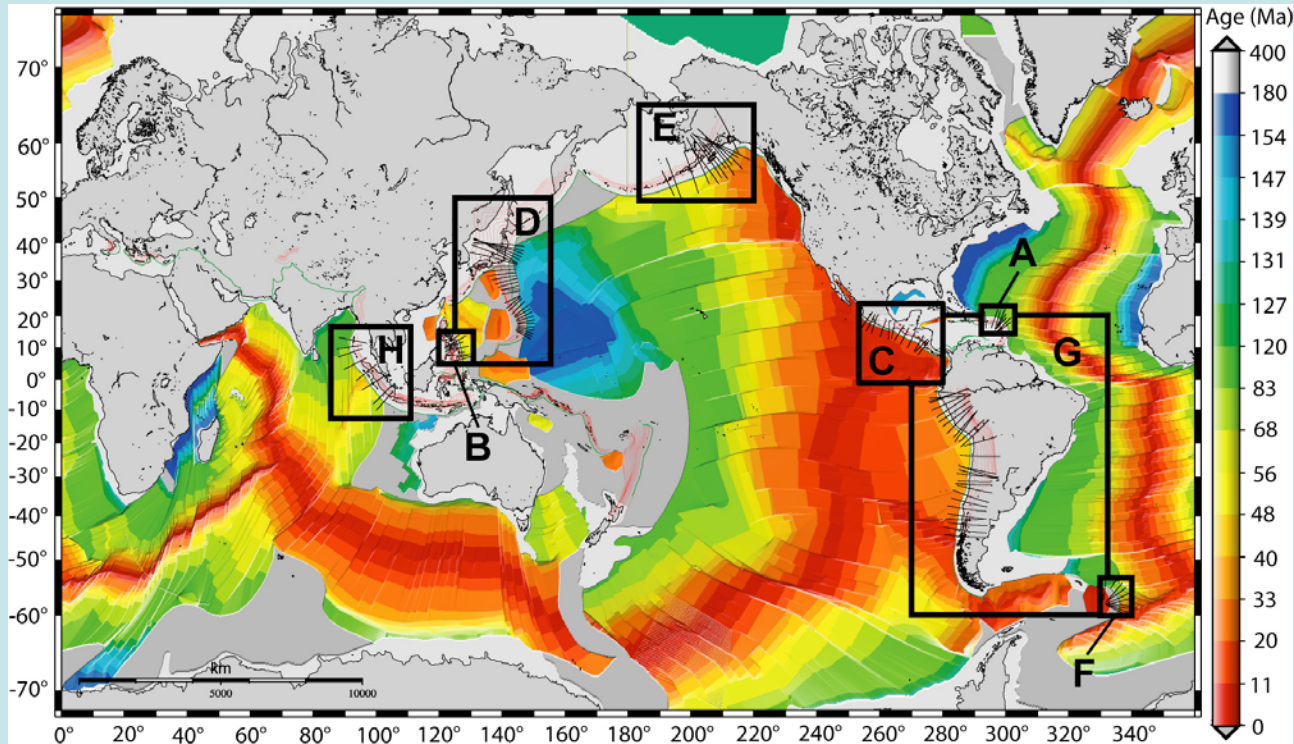
Plate margins rather move

Poligonal shapes of cells:

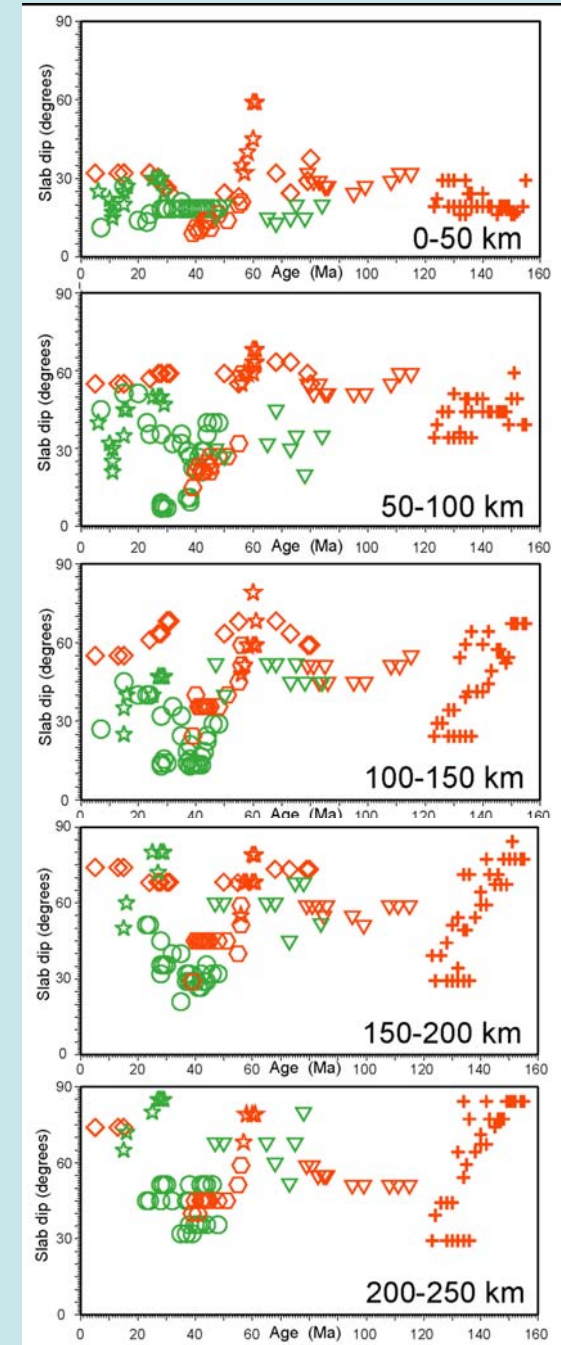
Plate margins can be very linear e.g., the Atlantic ridge



Tackley et al., 2003

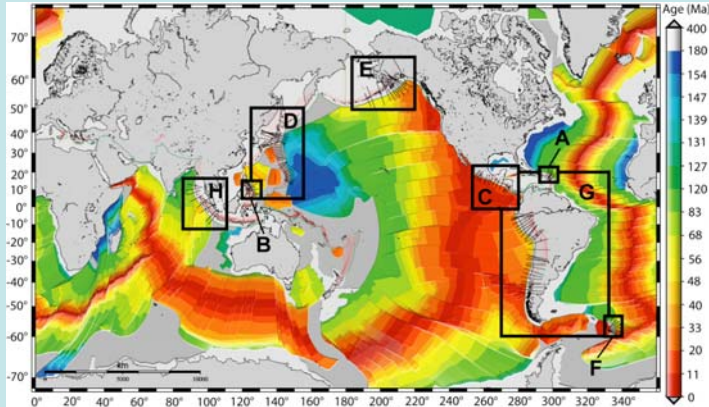


- South America
- ☆ Central America
- ▽ W-Indonesia
- ◇ E-Aleutians
- ▽ Caribbeans
- ☆ Philippines
- + Marianas - Japan

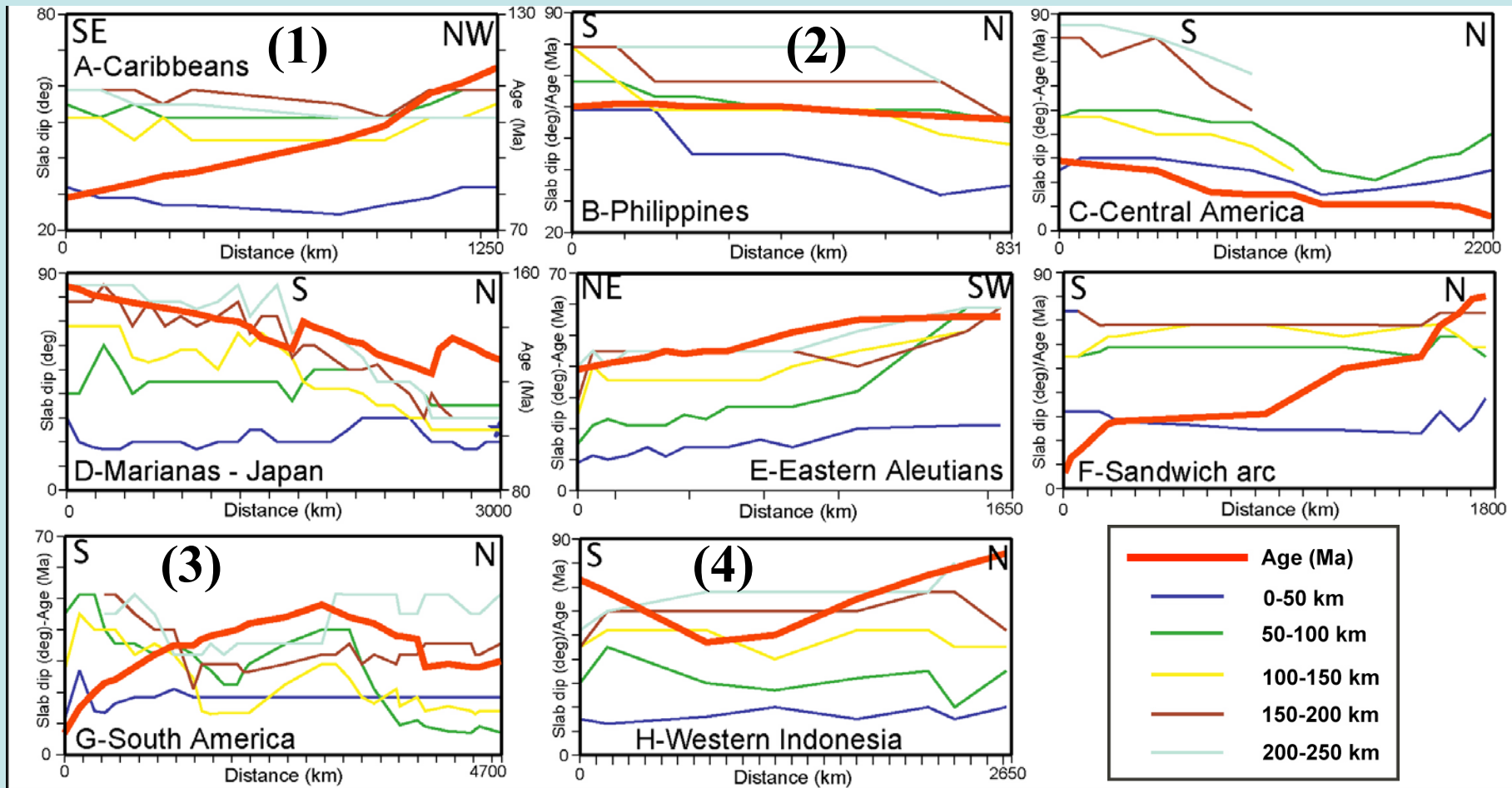


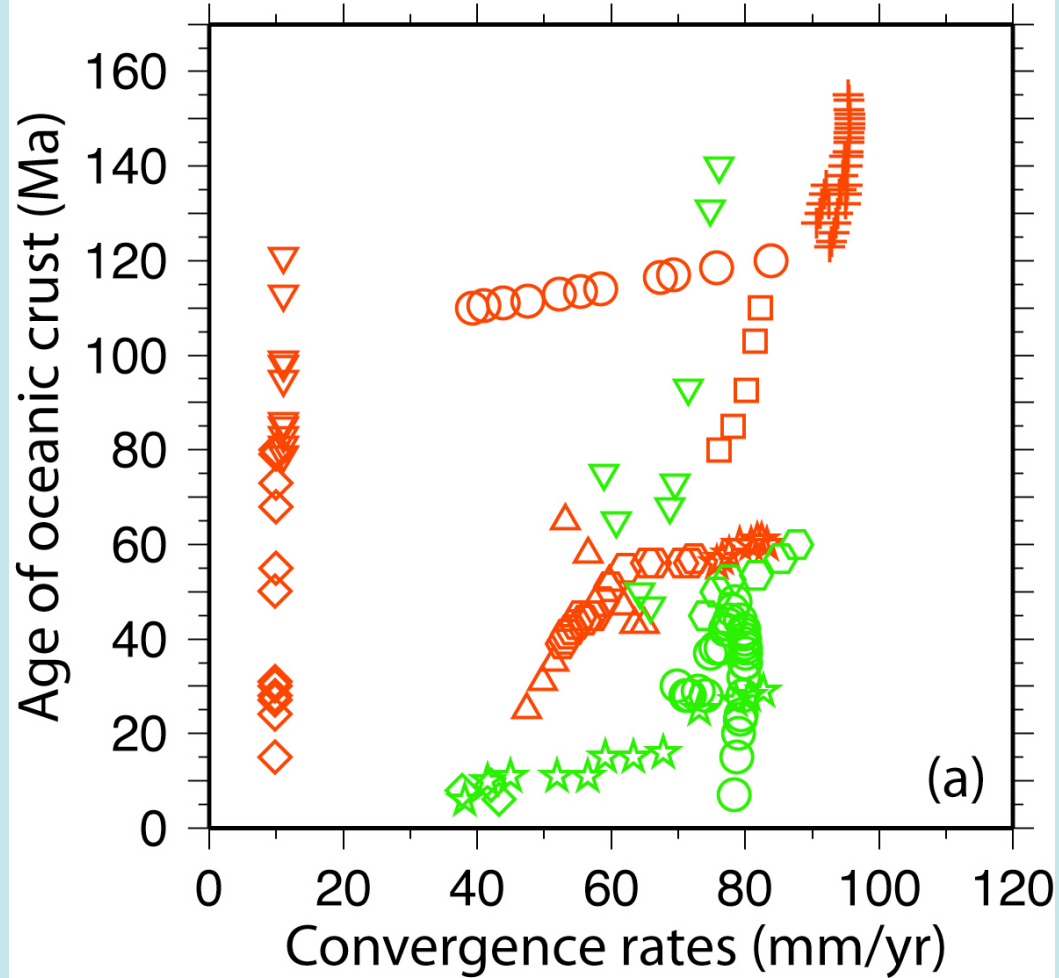
Slab dip Vs. Lithosphere age

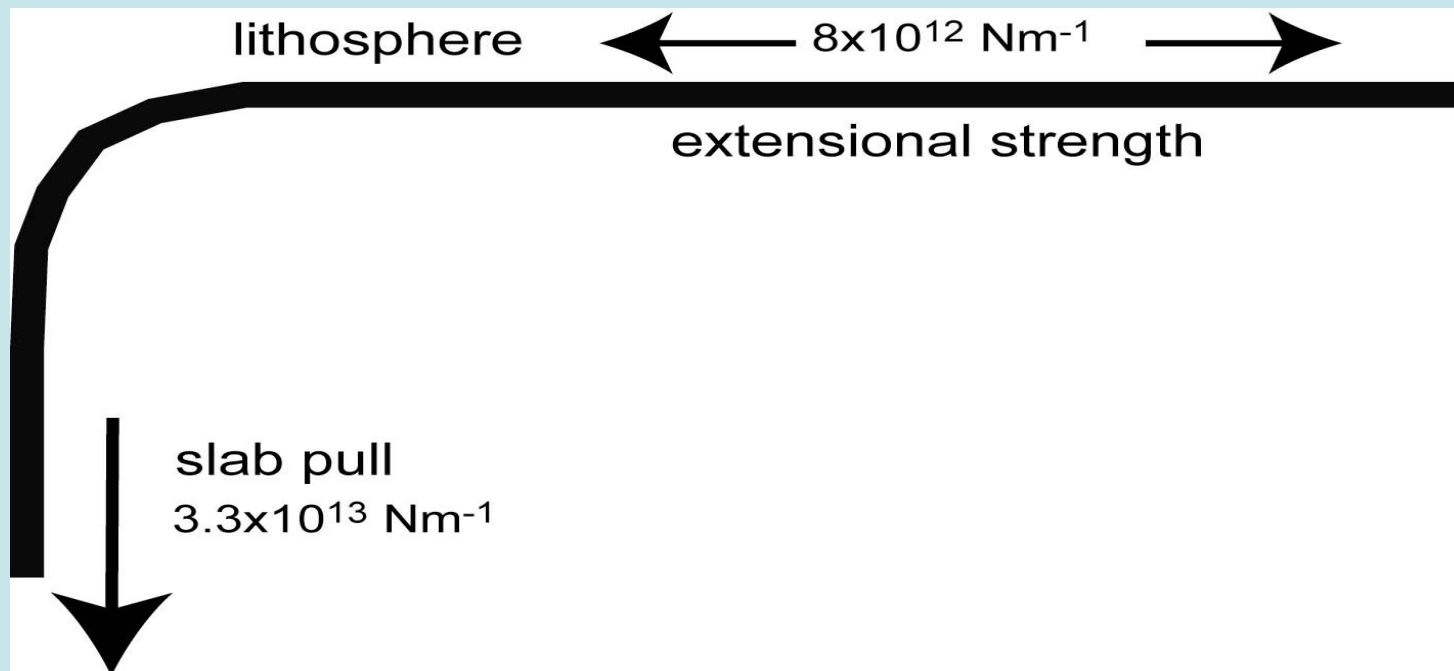
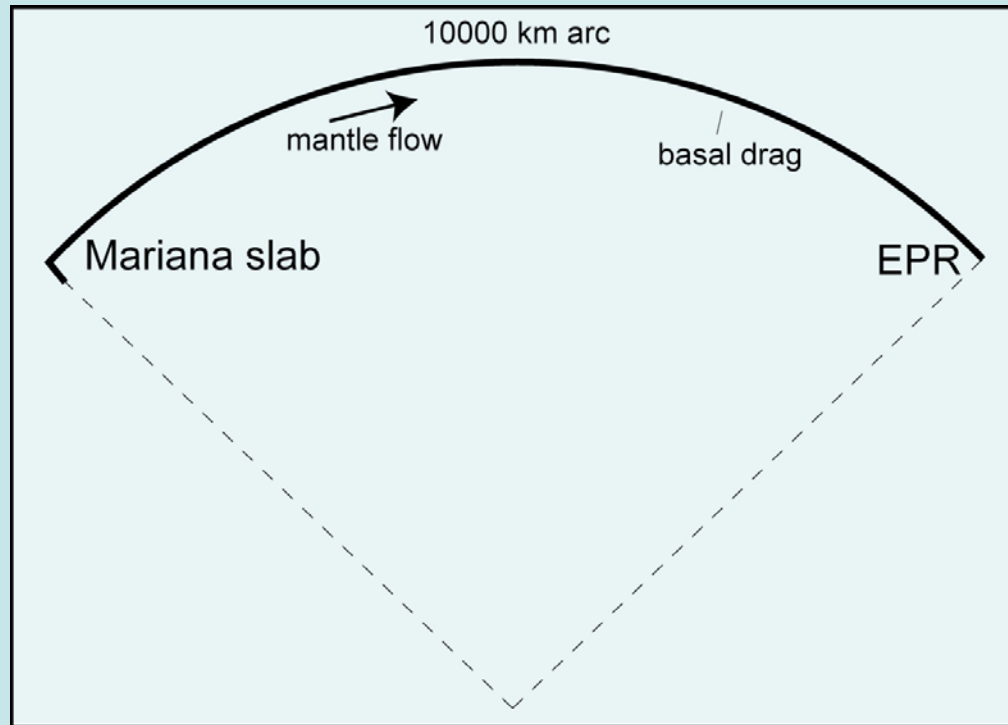
Same dip for variable ages = no correlation



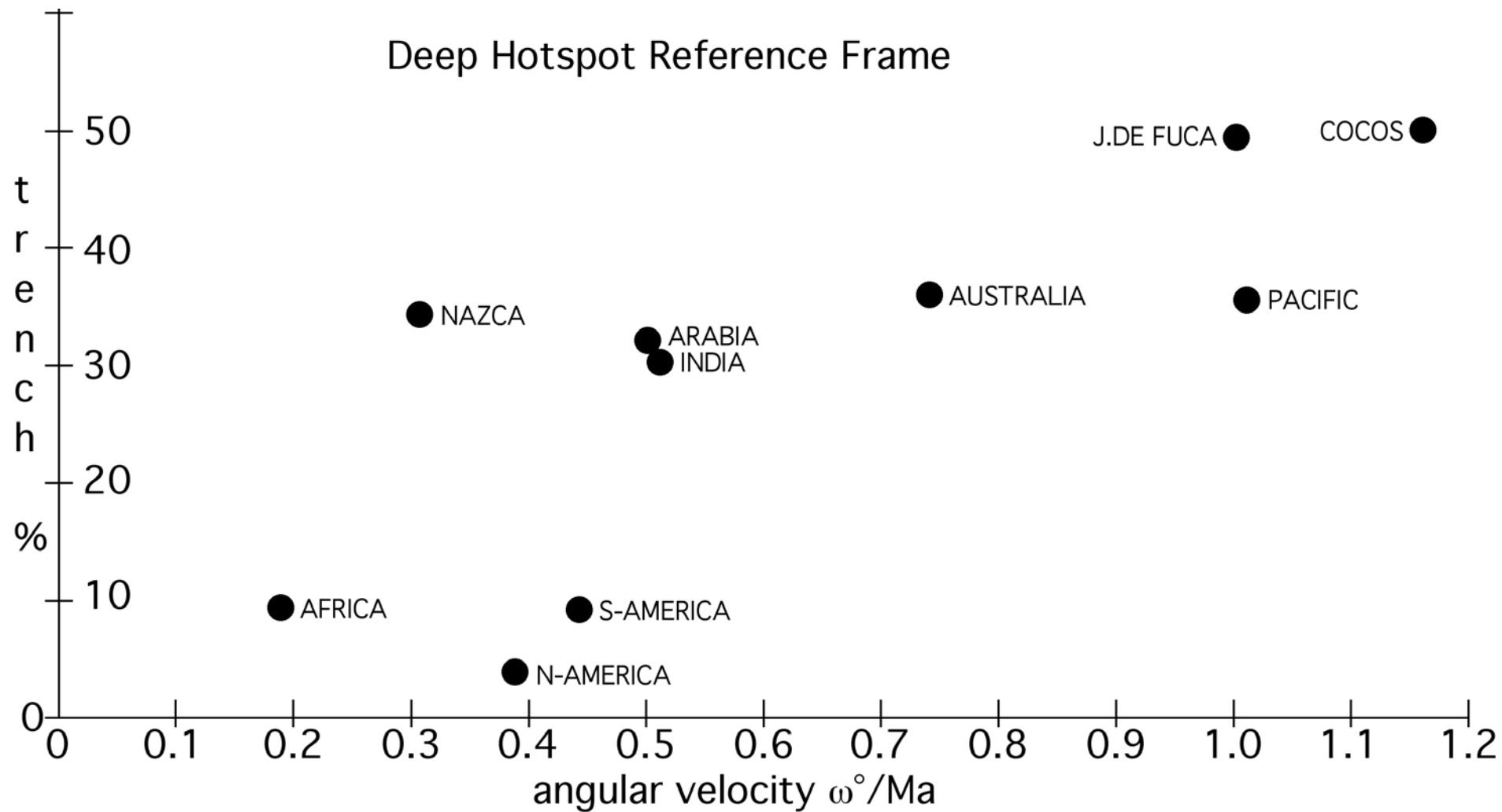
- Age increases but dip remains constant (1)
- Dip decreases but age remains constant (2)
- Dip decreases and age increases (3)
- Dip increases and age decreases (4)
- = No correlation



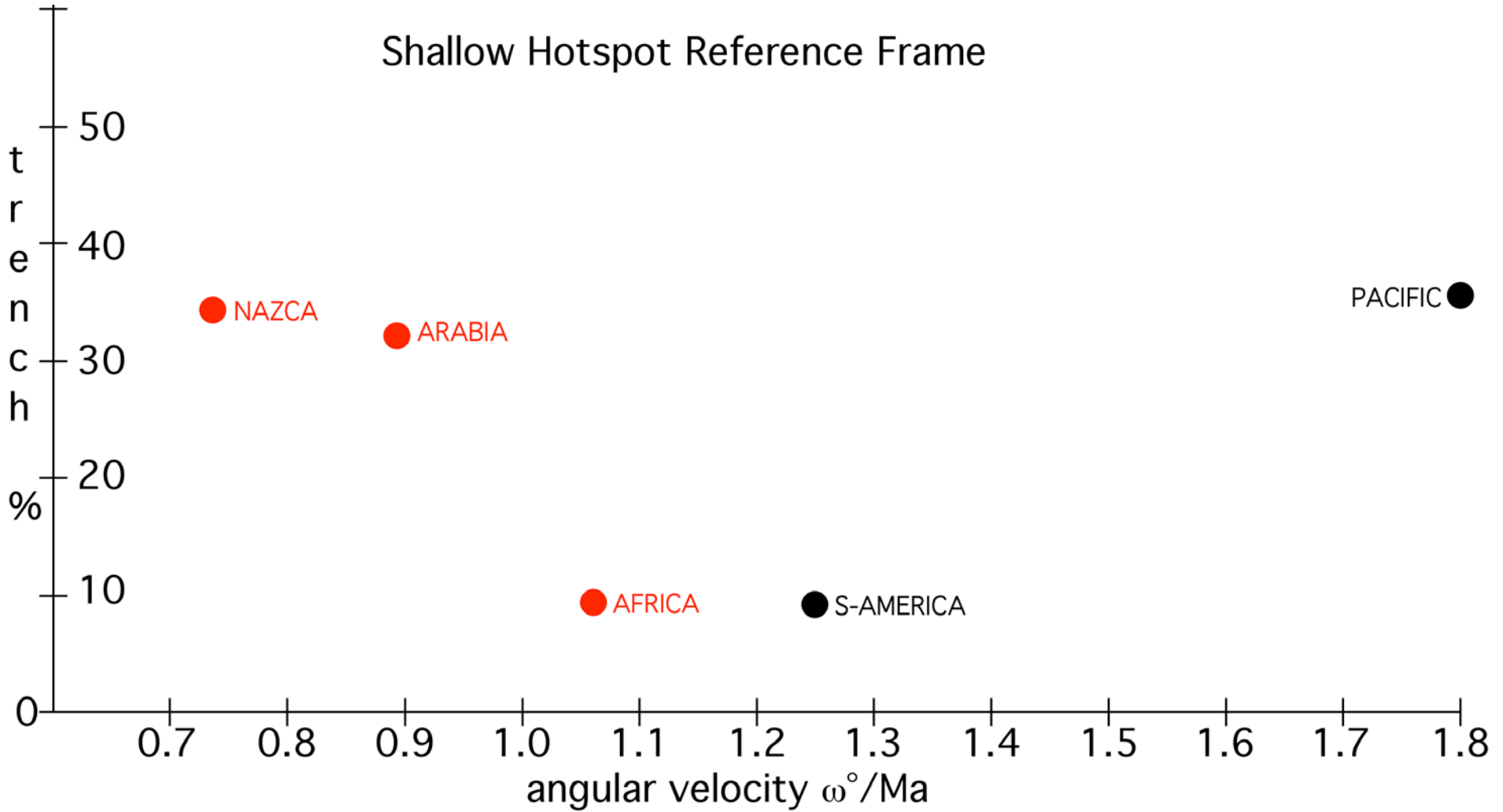


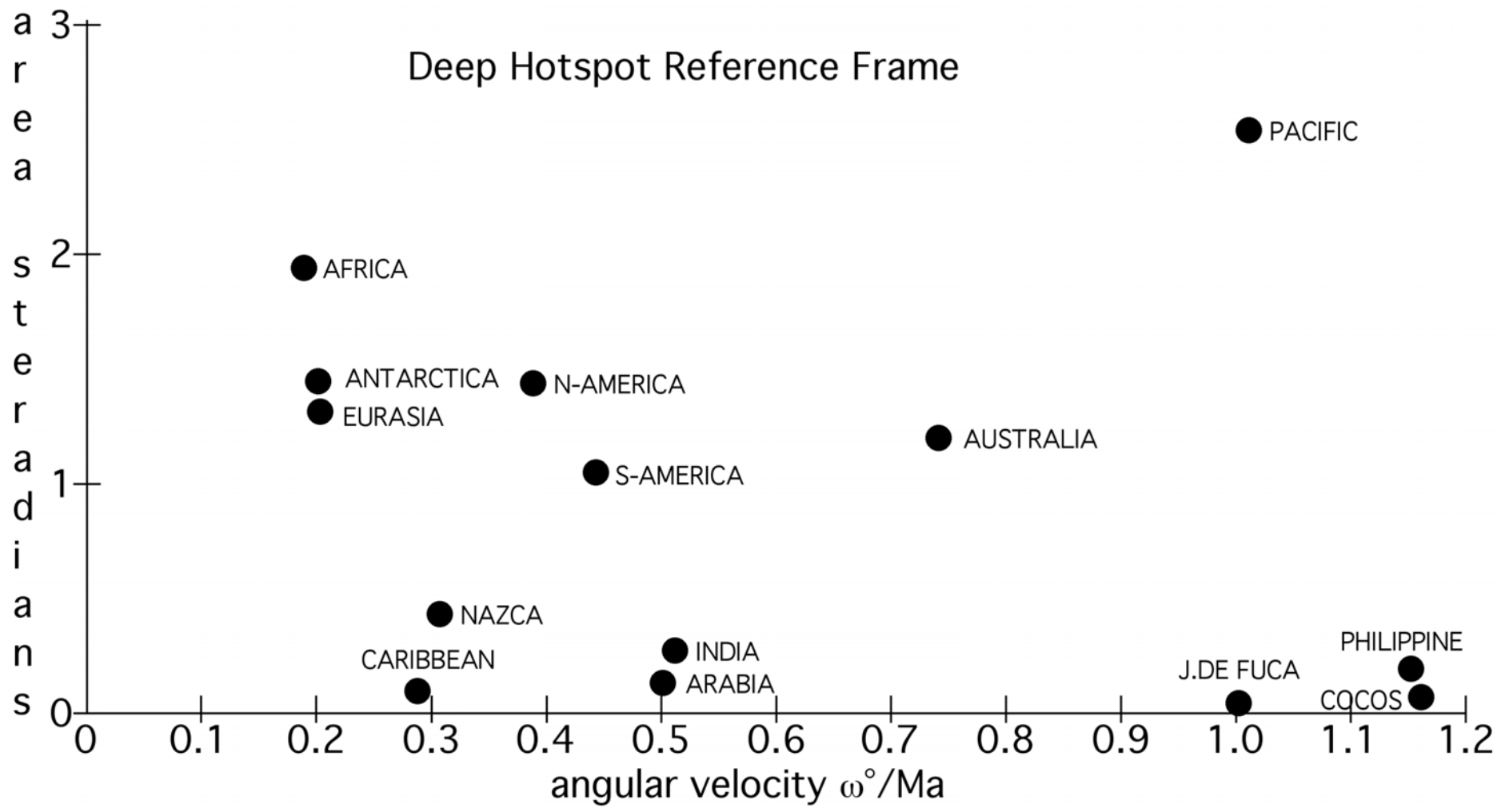


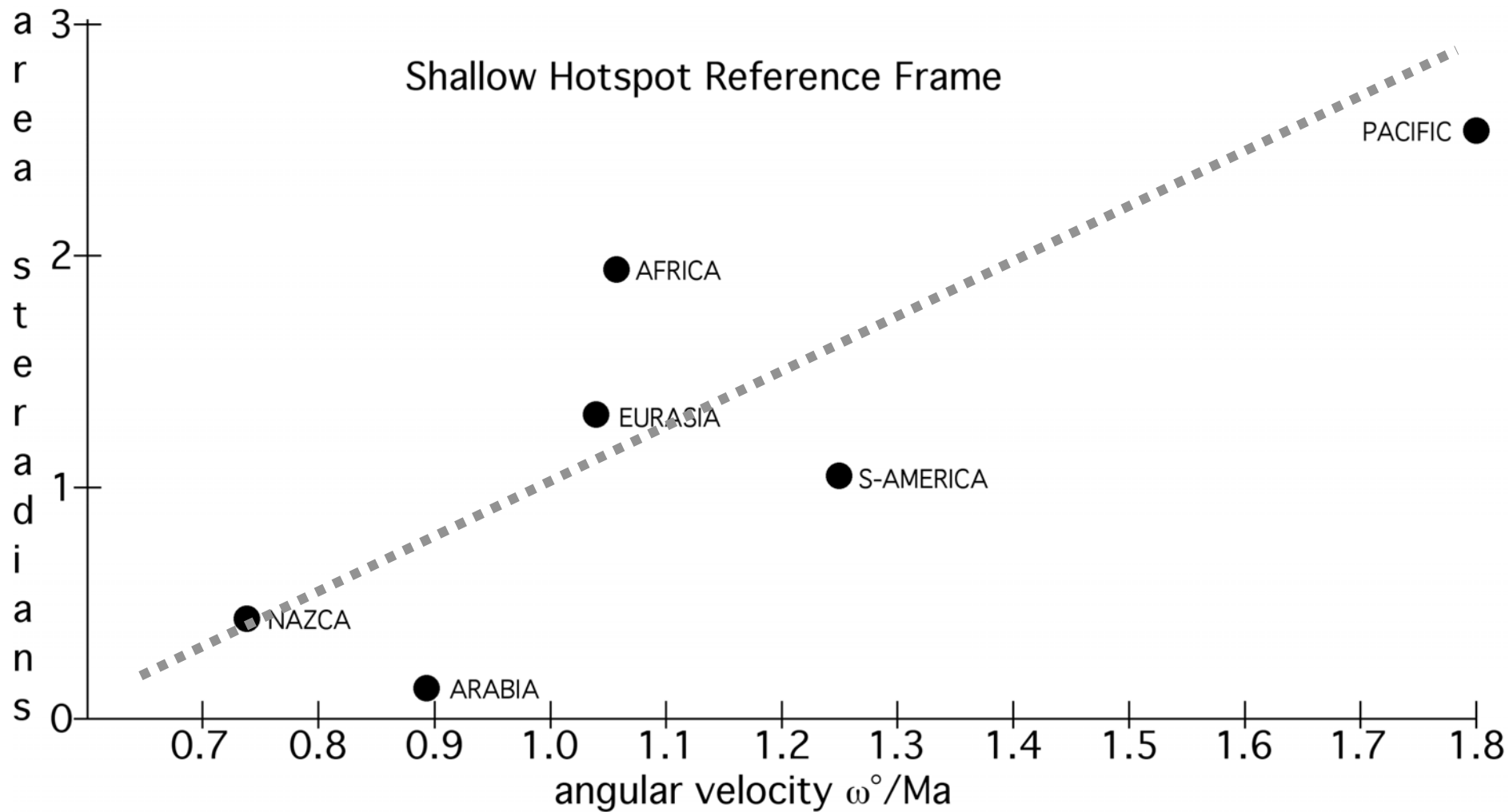
Deep Hotspot Reference Frame



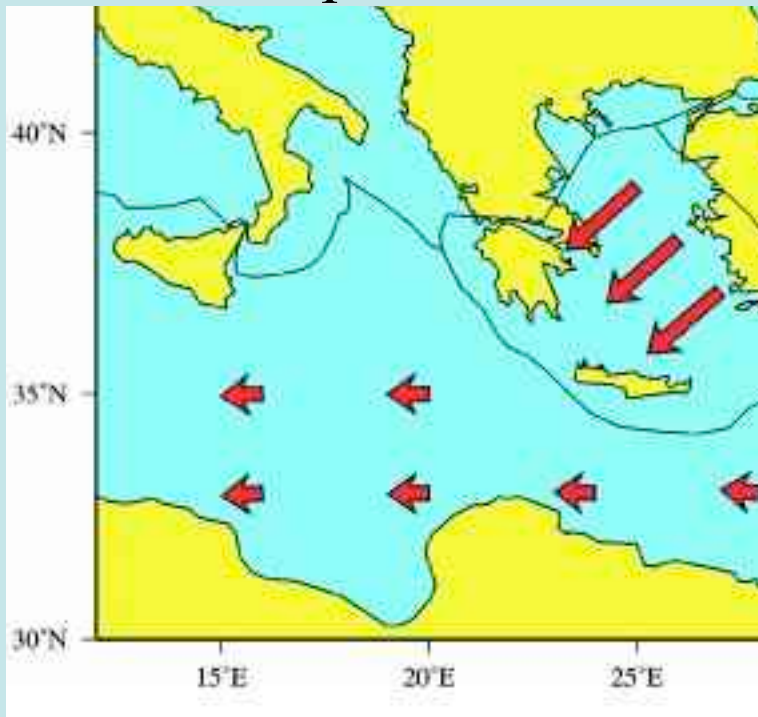
Shallow Hotspot Reference Frame





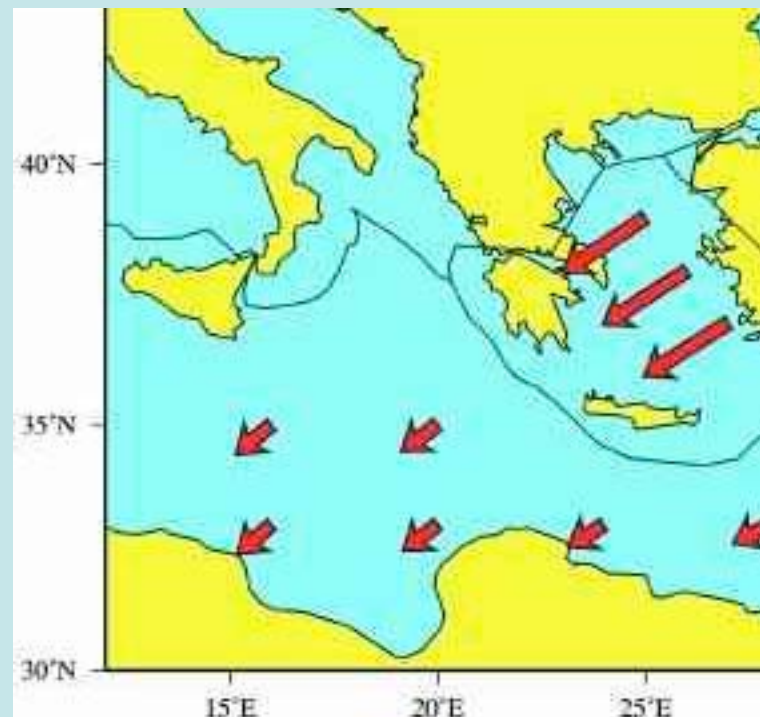


Deep HRF



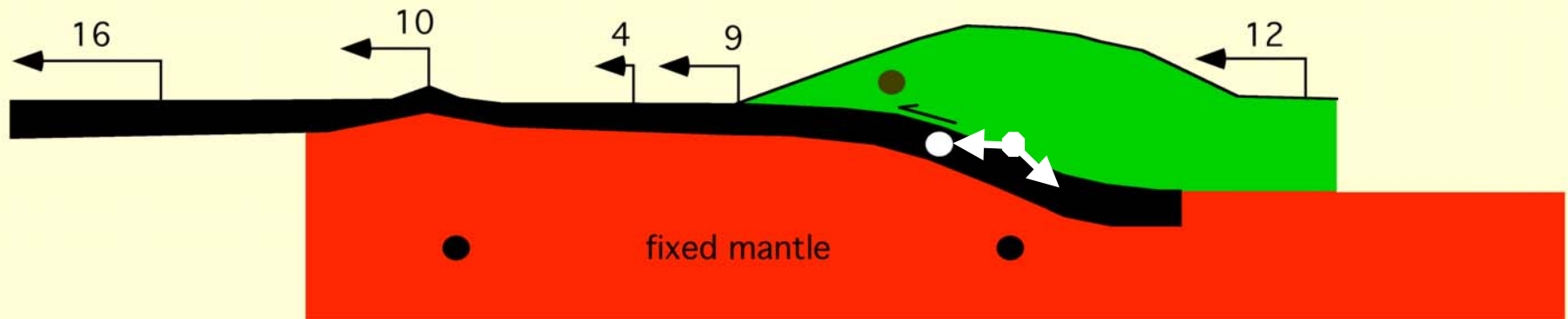
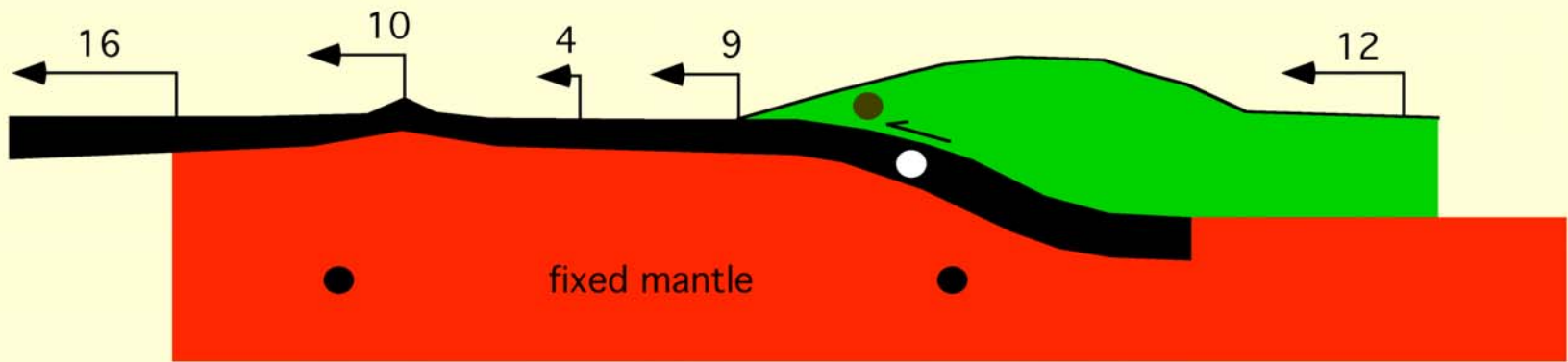
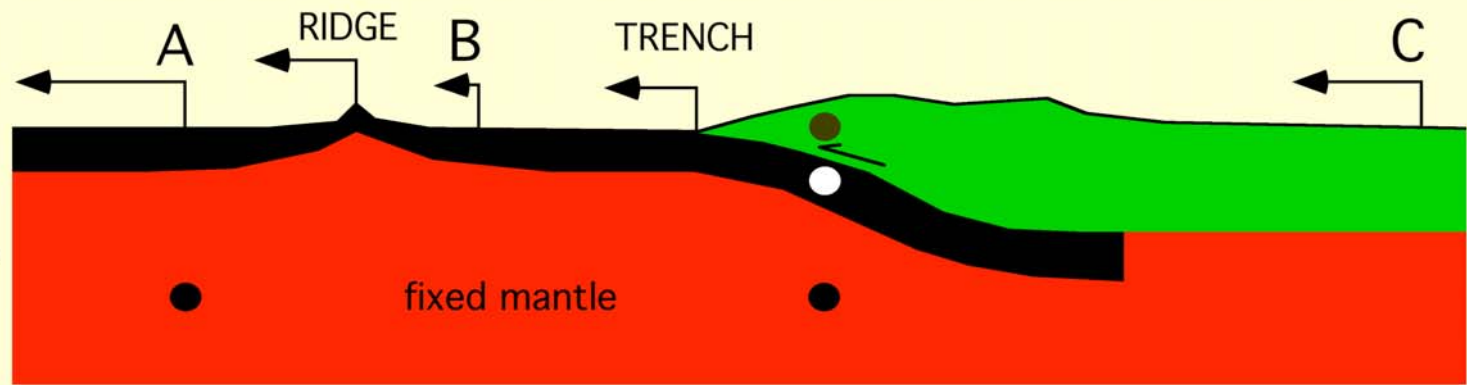
←
50 mm/yr

Shallow HRF



←
100 mm/yr

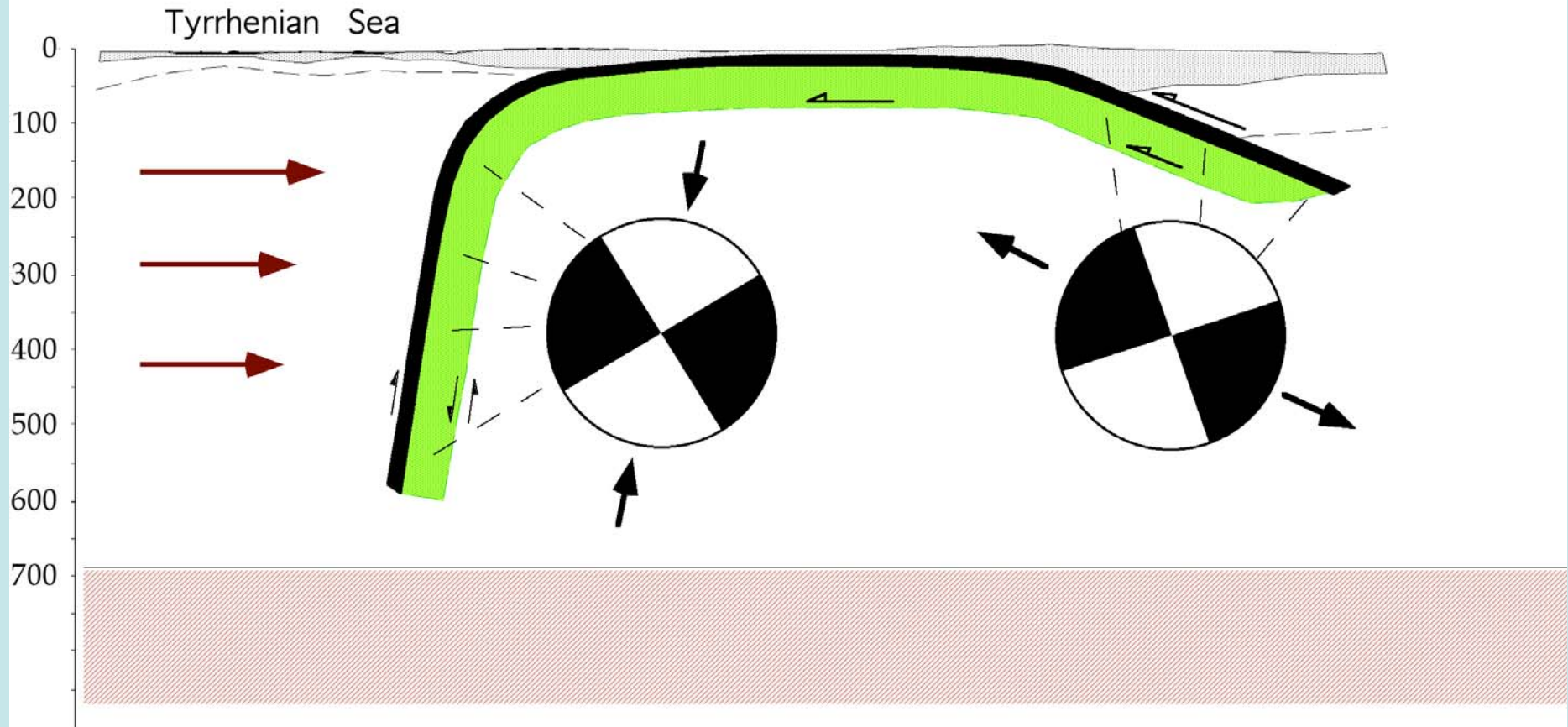
Plate A=16 cm/yr
 Plate B=4 cm/yr
 Plate C=12 cm/yr
 convergence=8 cm/yr
 subduction=5 cm/yr
 shortening=3 cm/yr



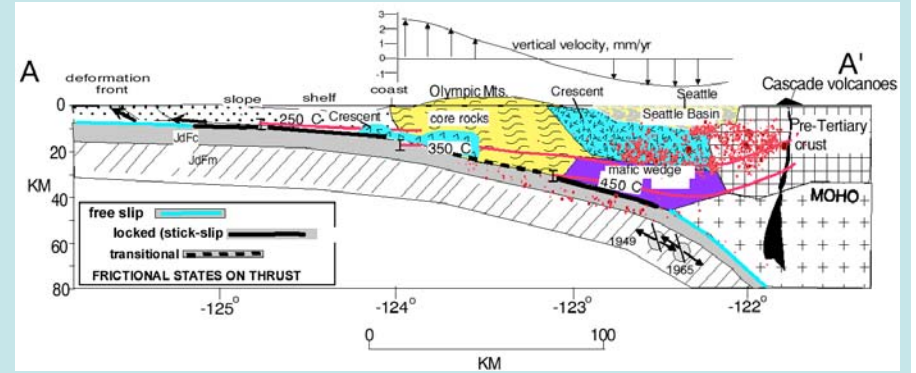
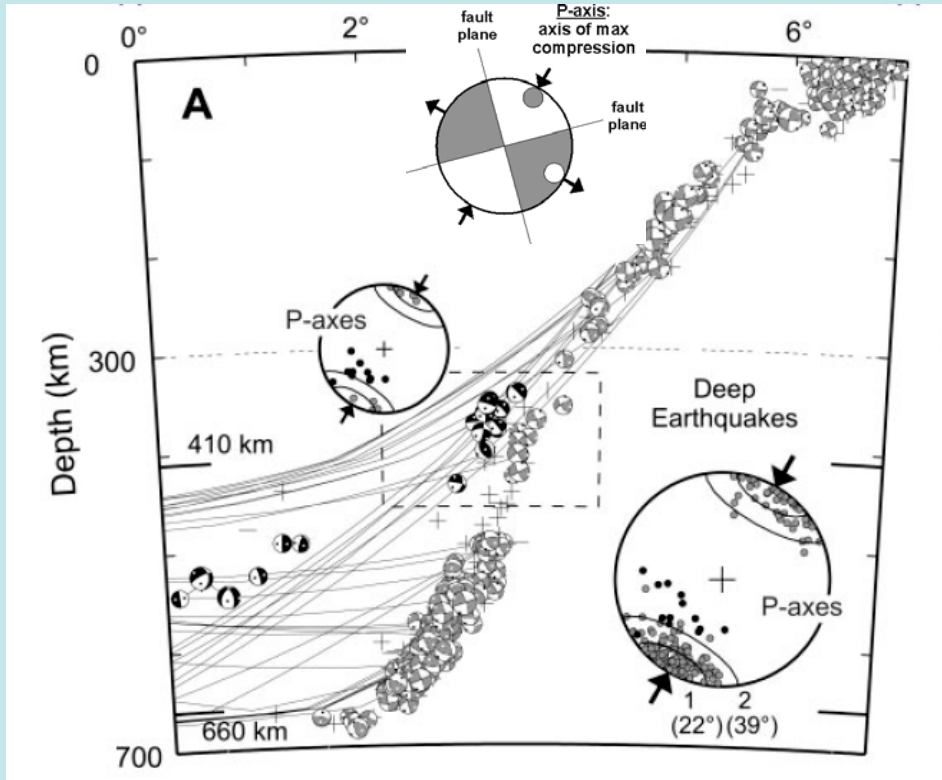
slab remoting west at 4 cm/yr relative to mantle
 upper plate overriding the lower plate at 5 cm/yr

DOWN-DIP COMPRESSION
Apennines slab

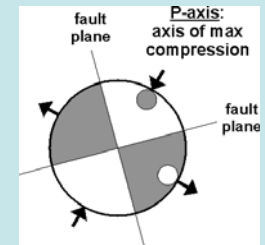
DOWN-DIP EXTENSION
Hellenides slab



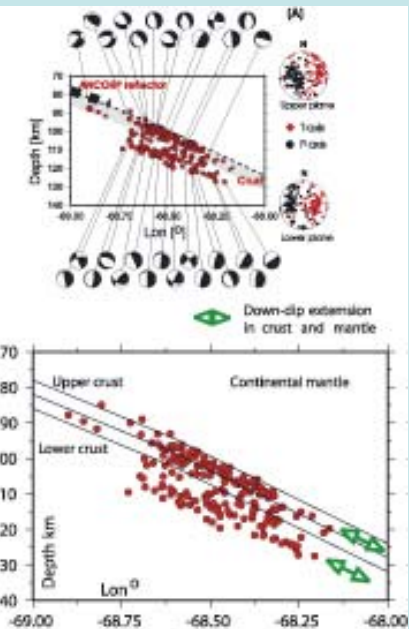
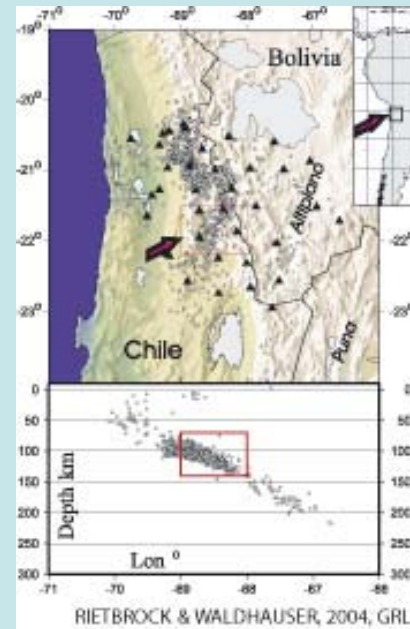
TONGA

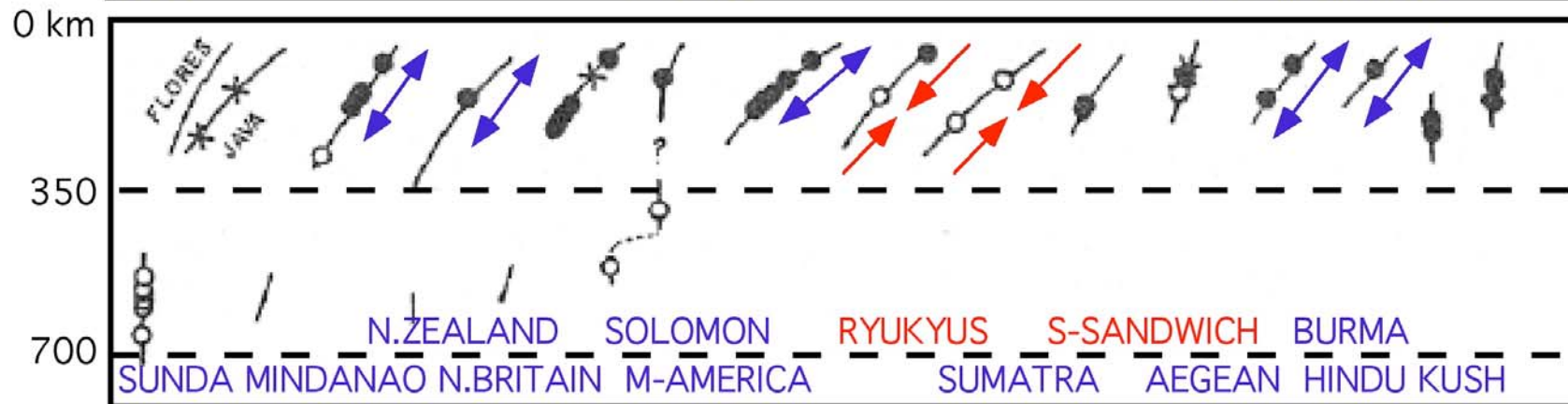
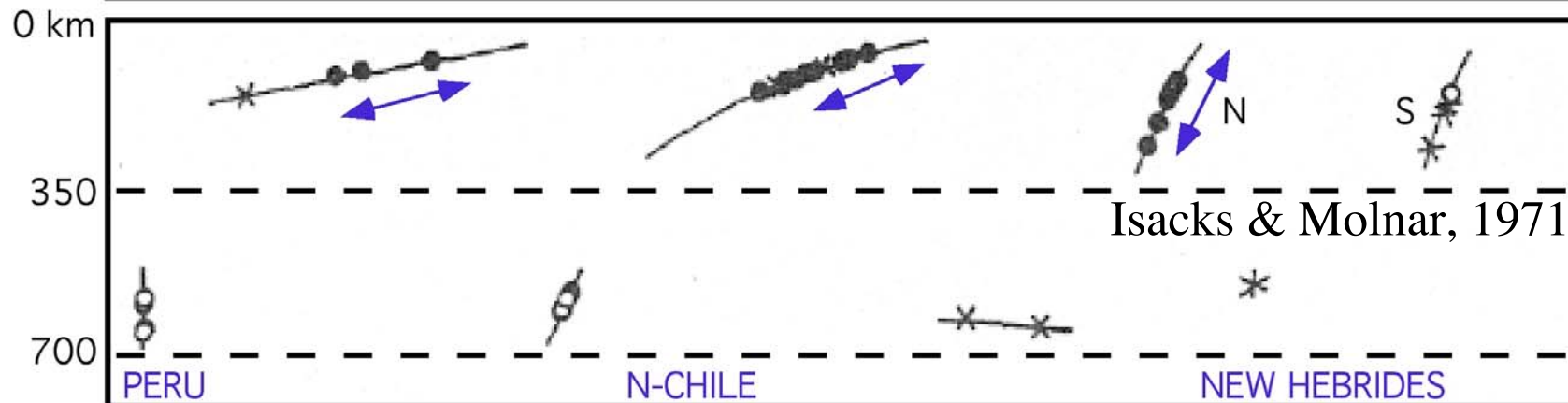
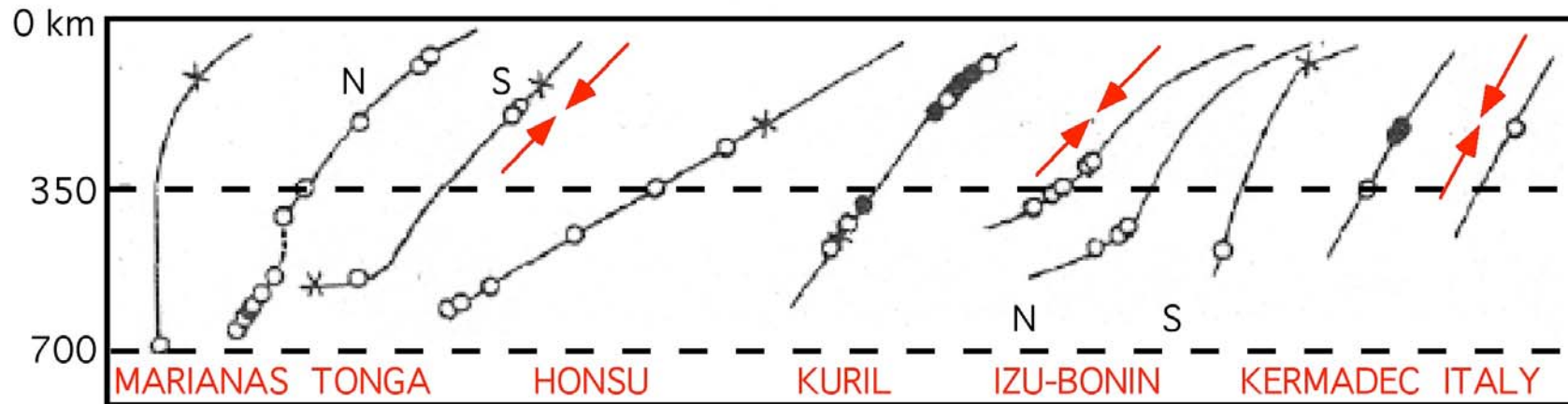


CASCADIA



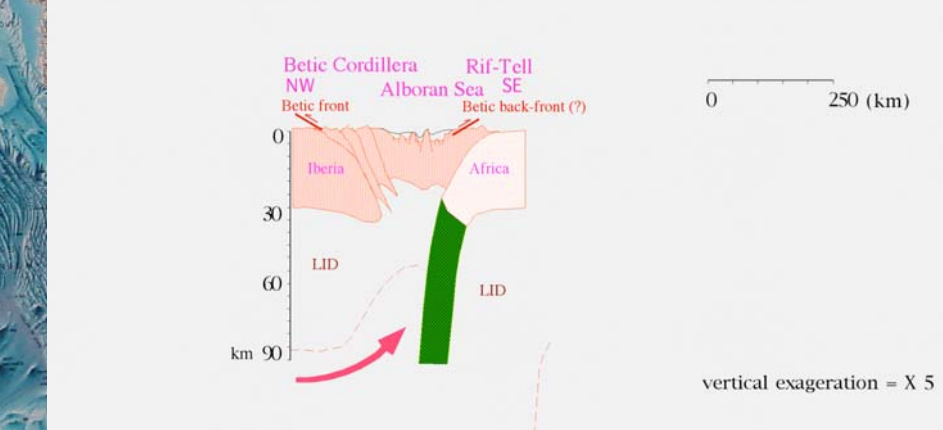
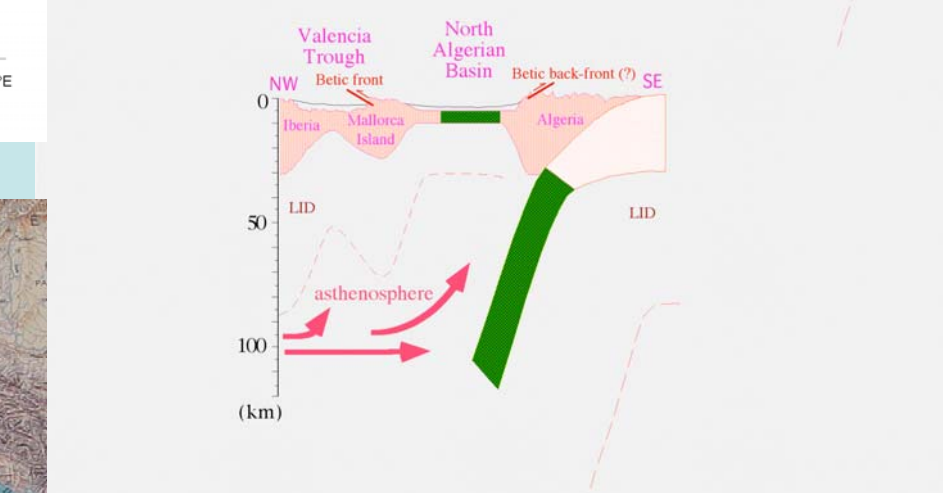
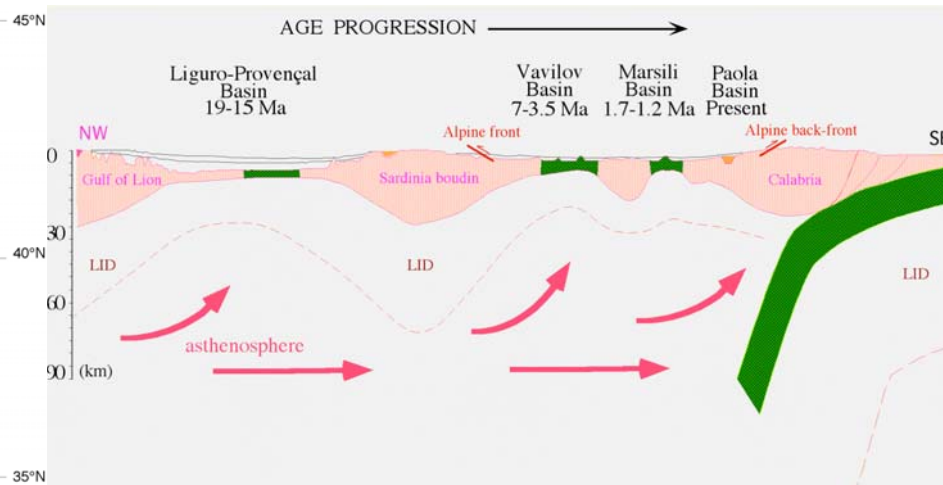
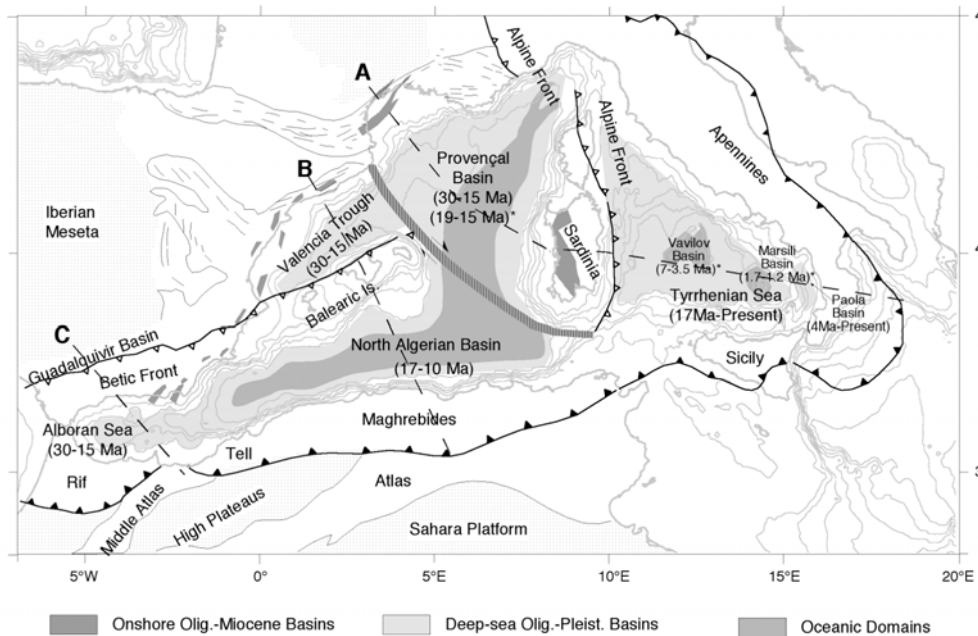
CHILE





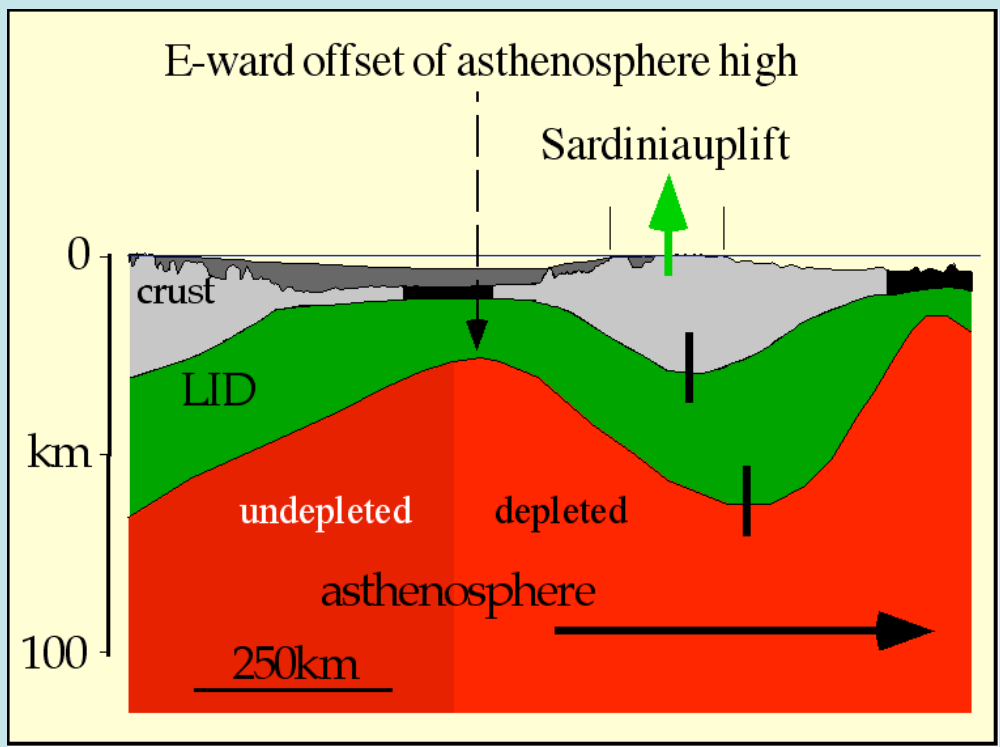
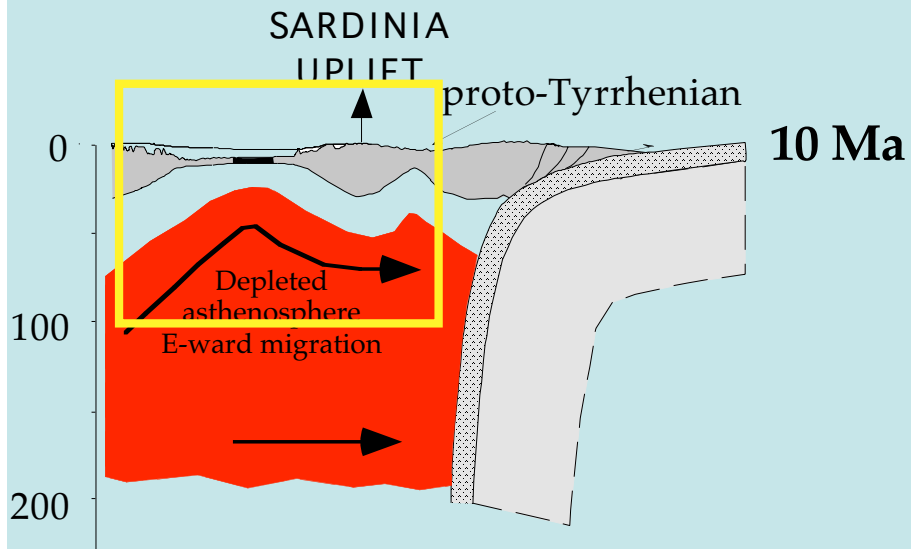
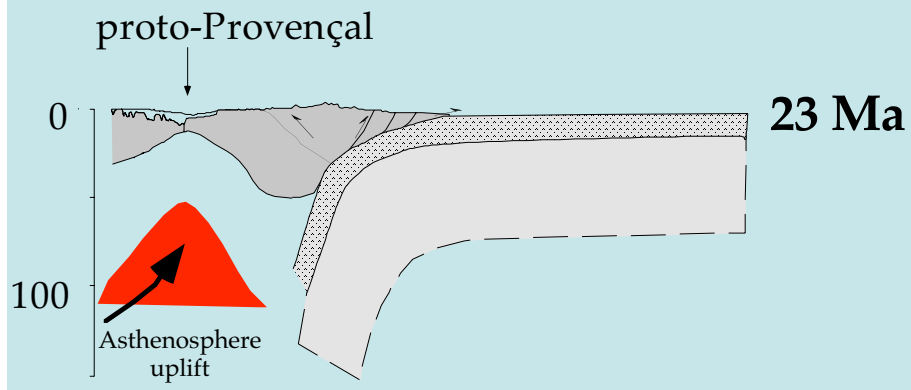
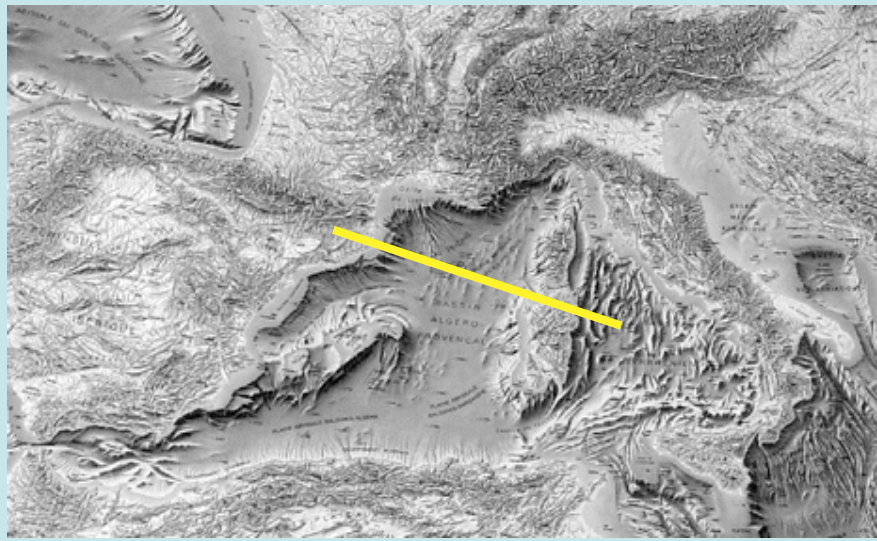
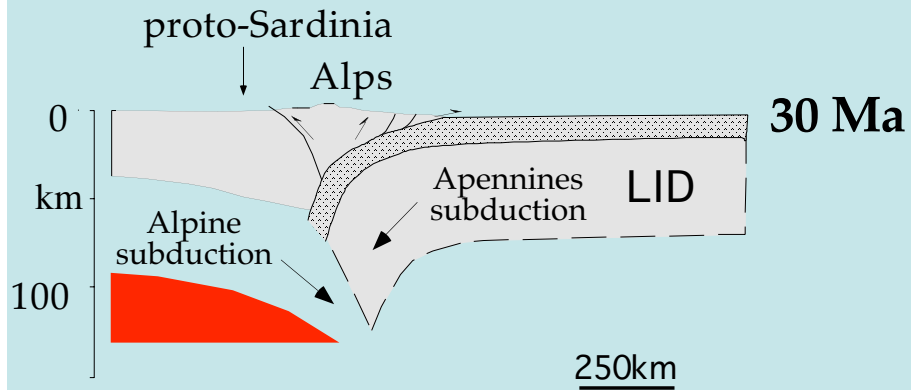
W-directed

E-NE-directed

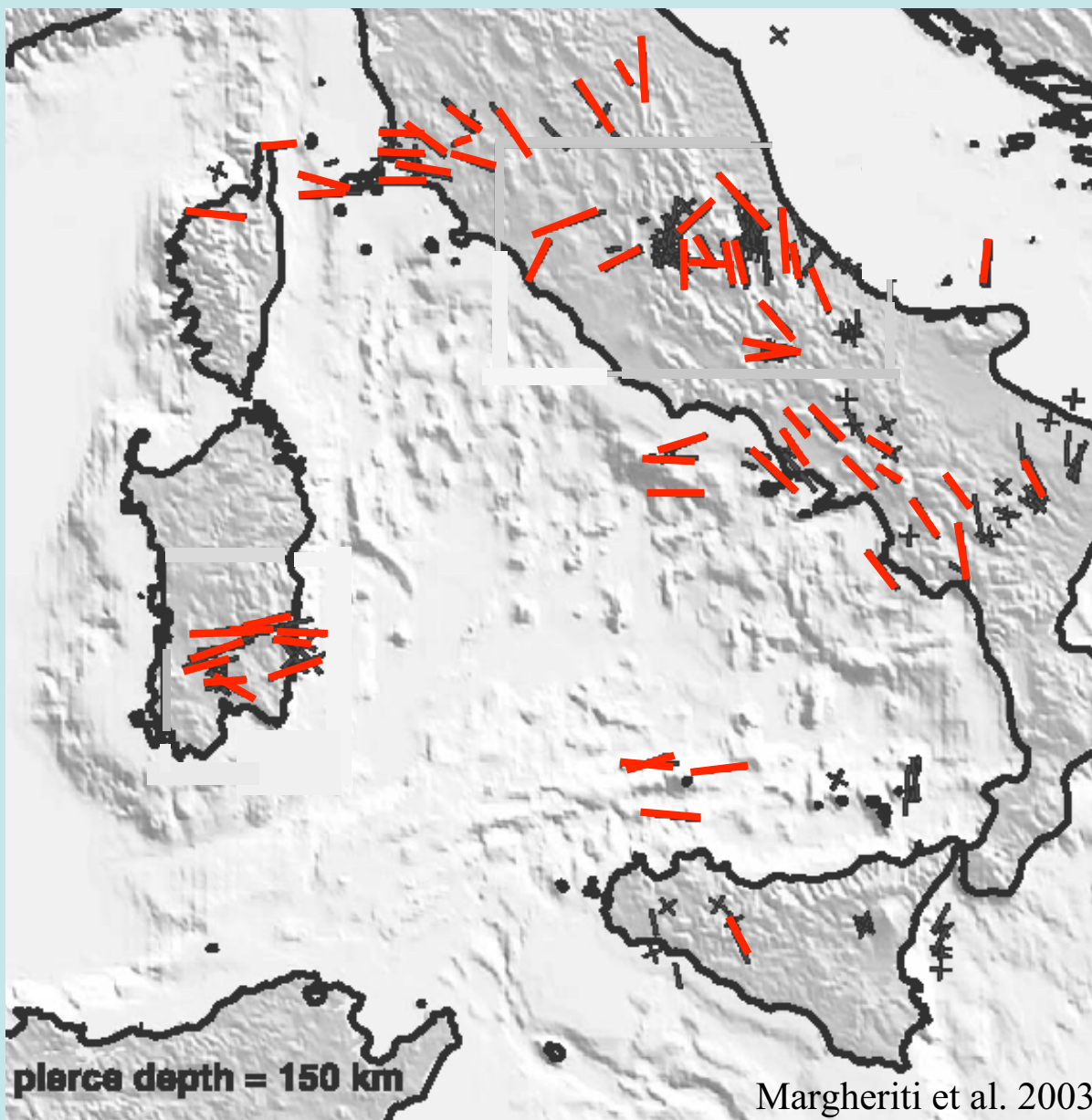
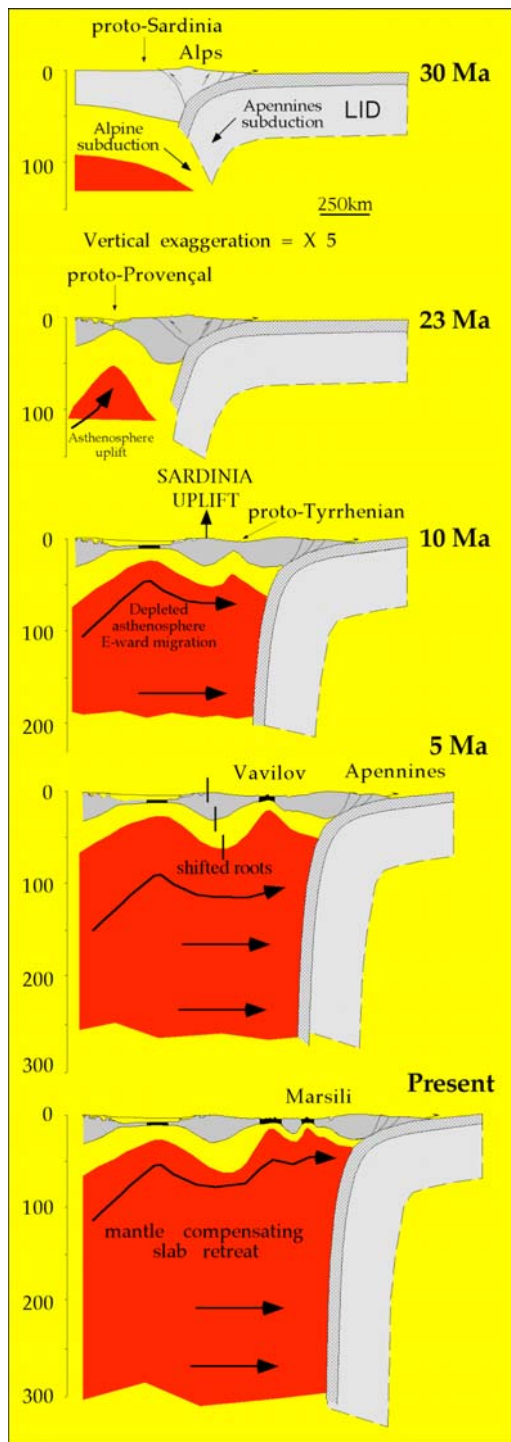


0 250 (km)

vertical exaggeration = X 5



Shear Wave Splitting



Margheriti et al. 2003

Mantle - ρ deficit



extension



Tyrrhenian coast line

extension

1980, Irpinia

1990, Potenza

compression

Adriatic coast line

Mantle wedging



Moho

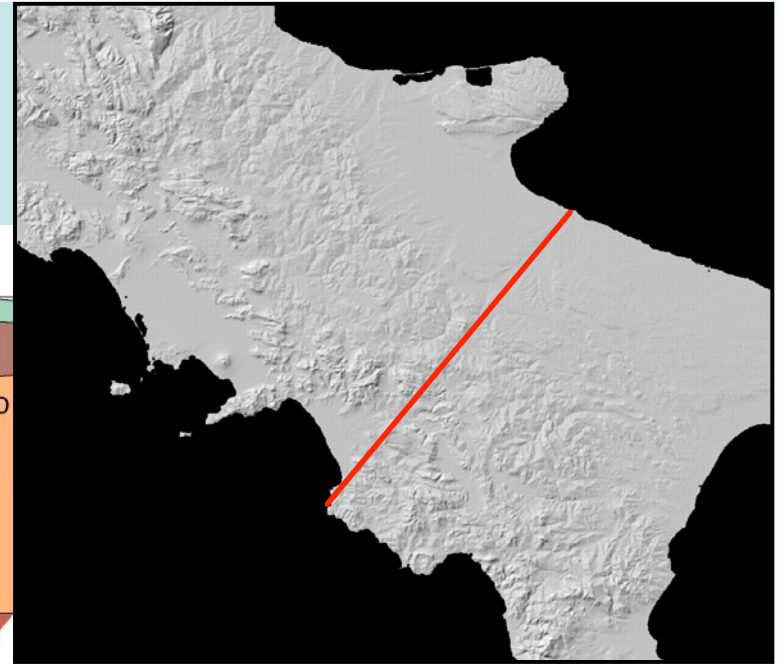
LID

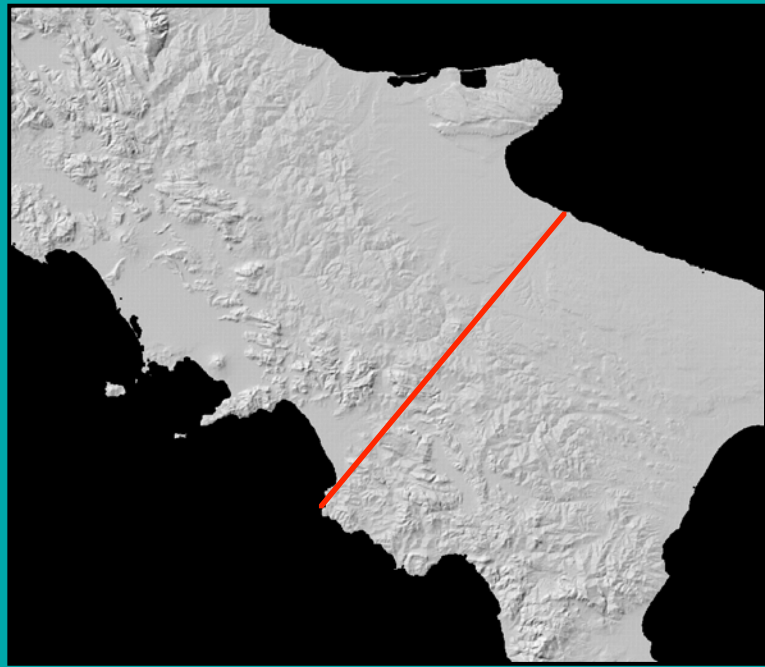
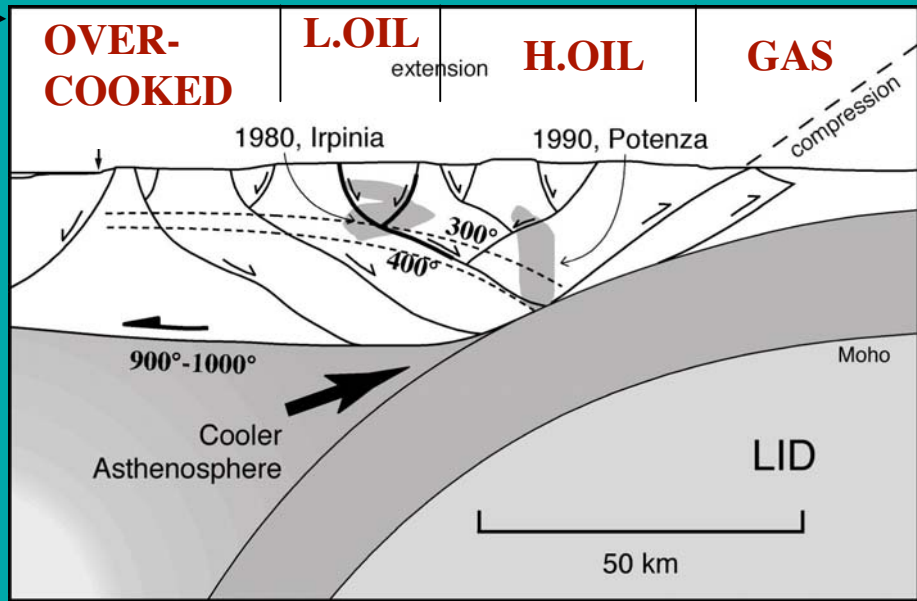
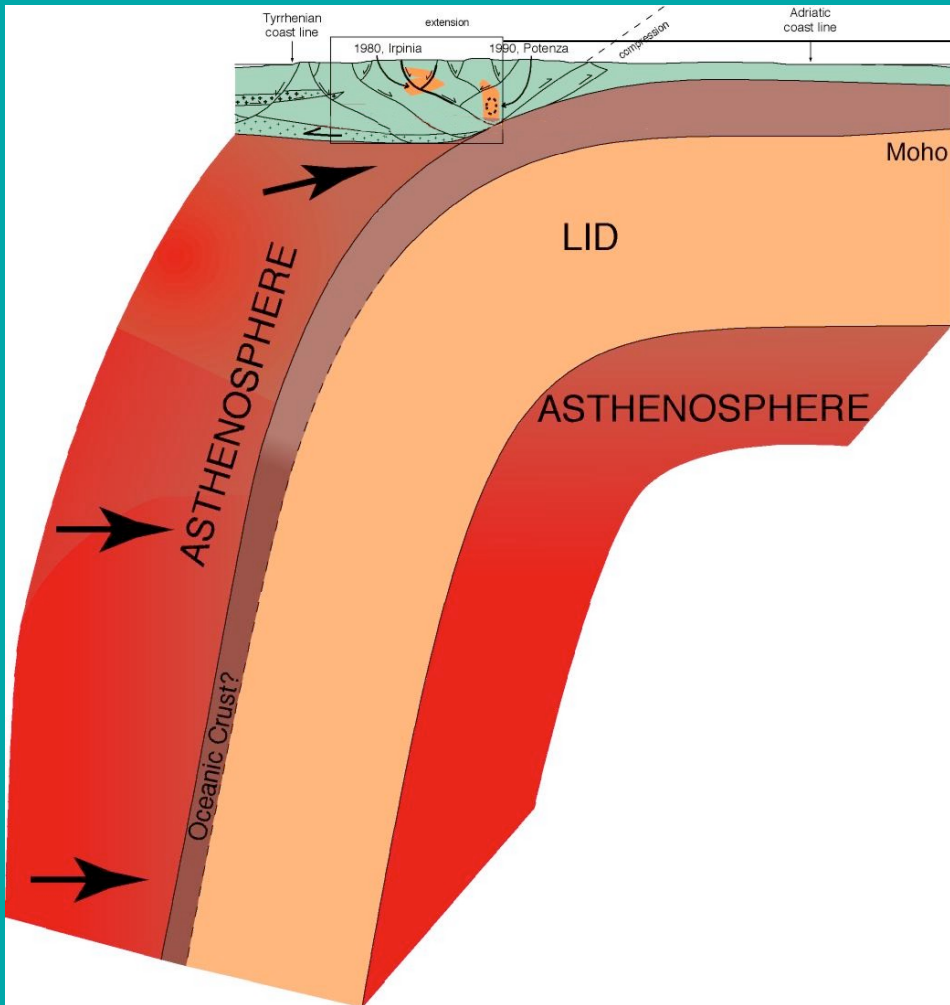
ASTHENOSPHERE

ASTHENOSPHERE

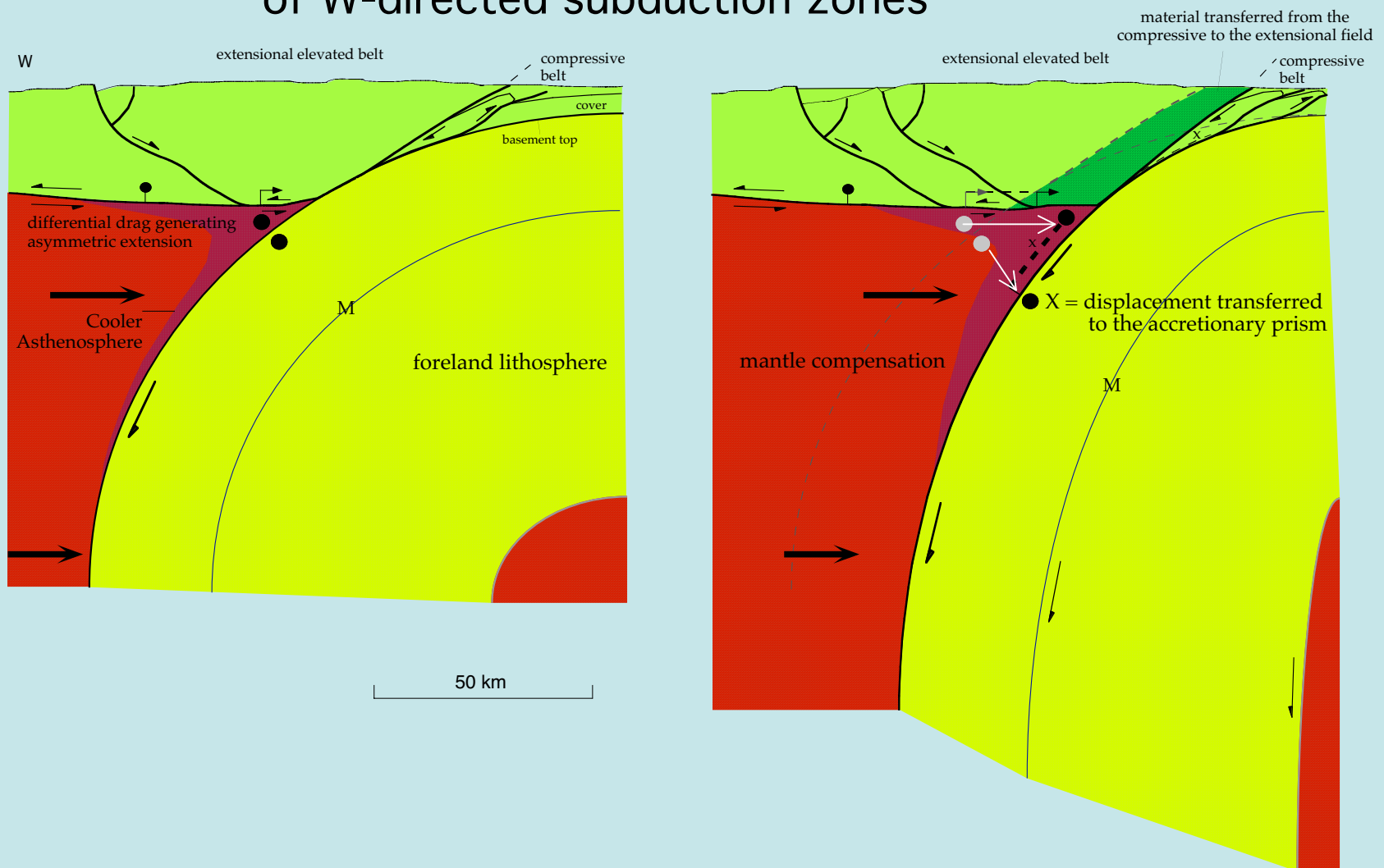


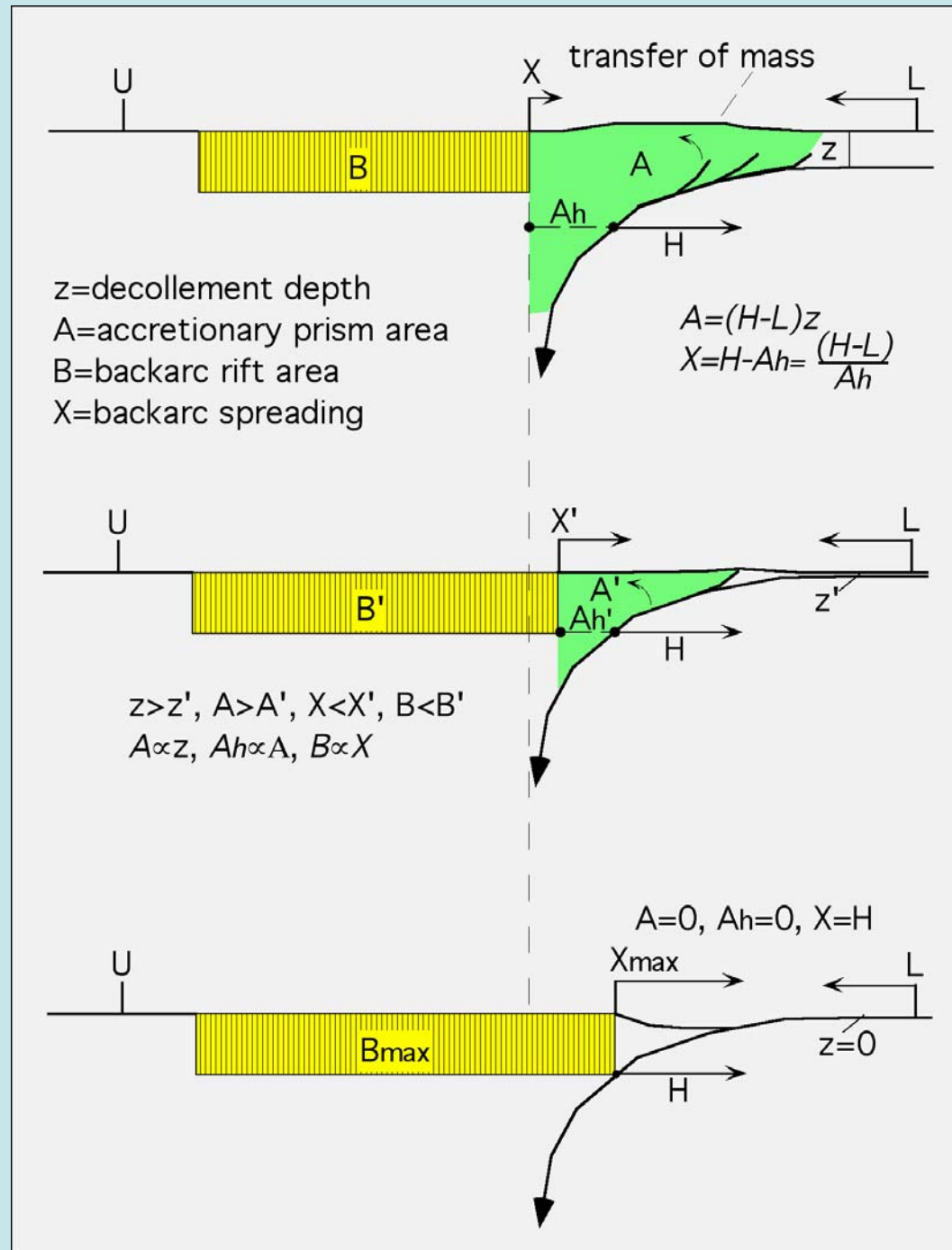
Oceanic Crust?

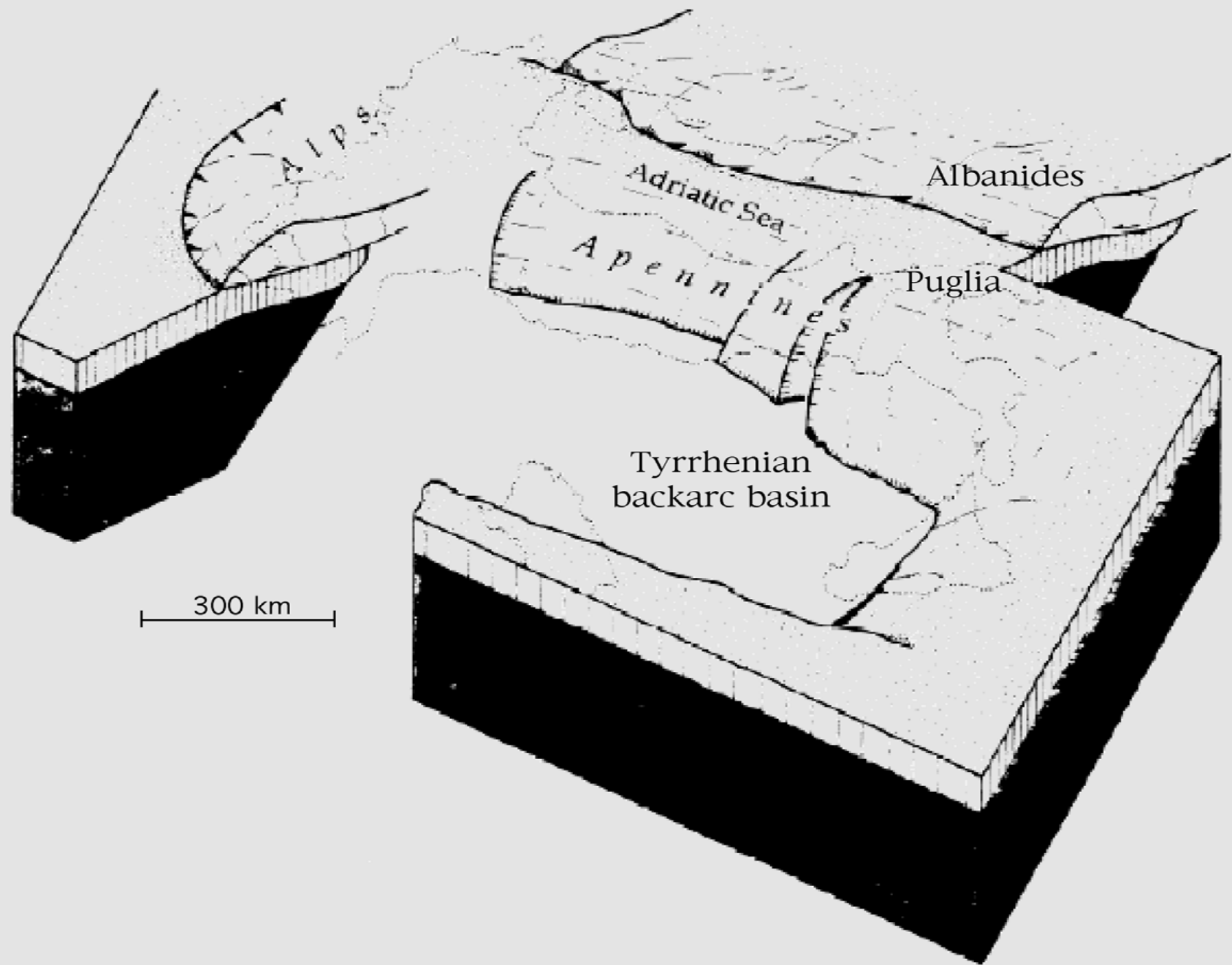


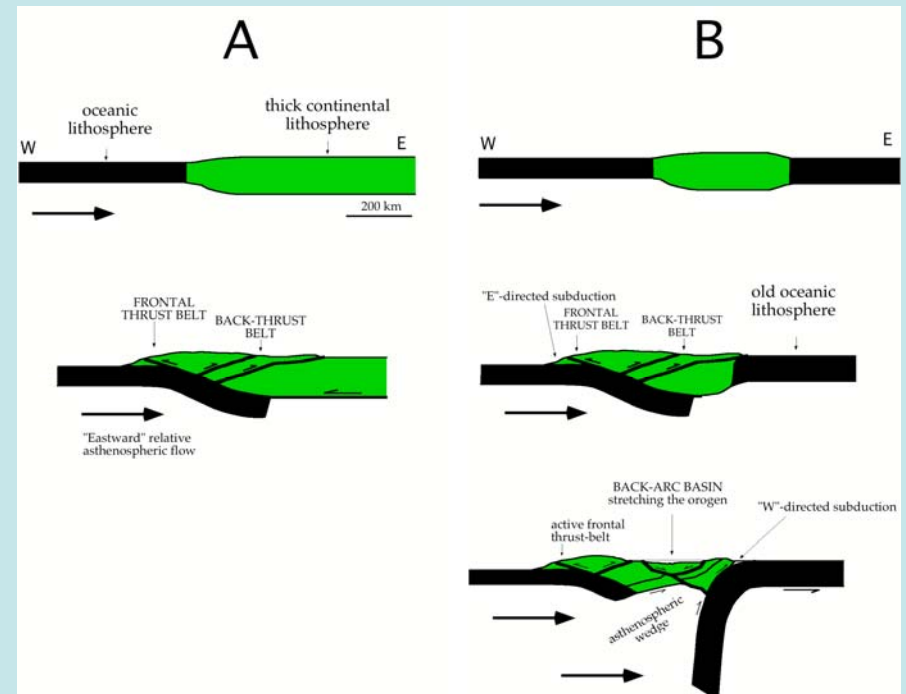
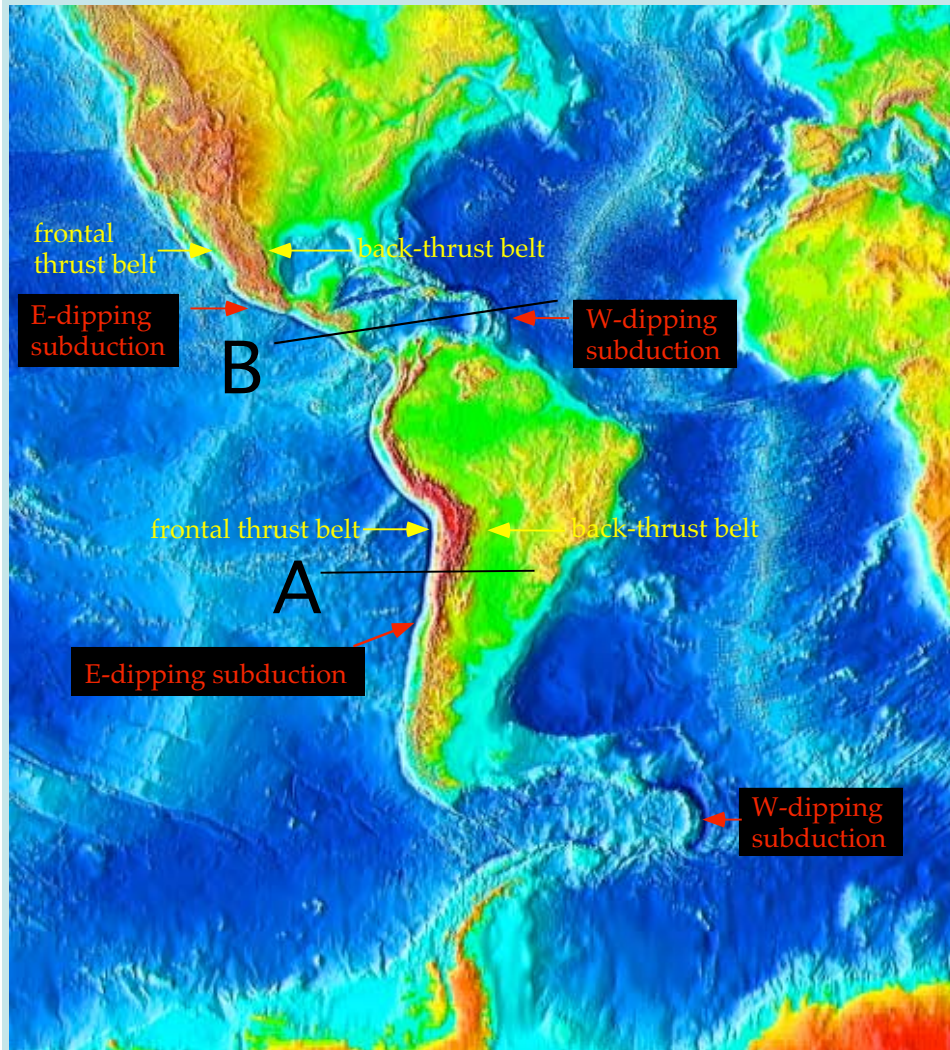


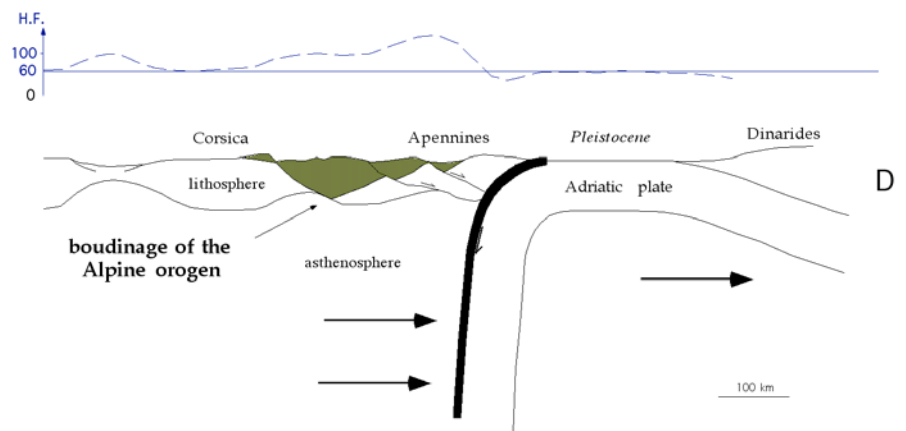
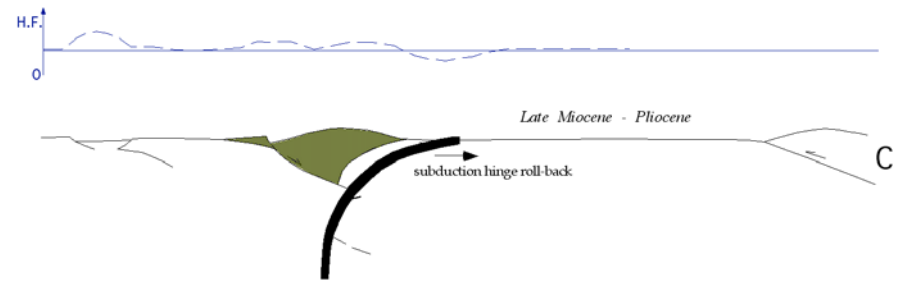
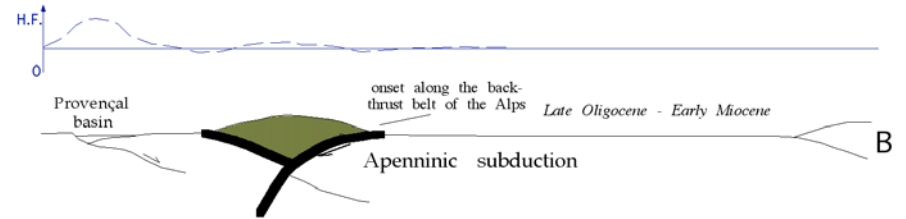
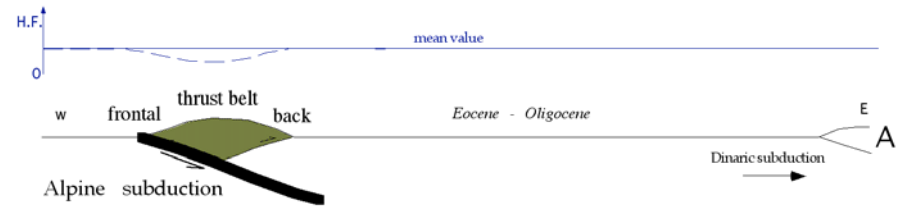
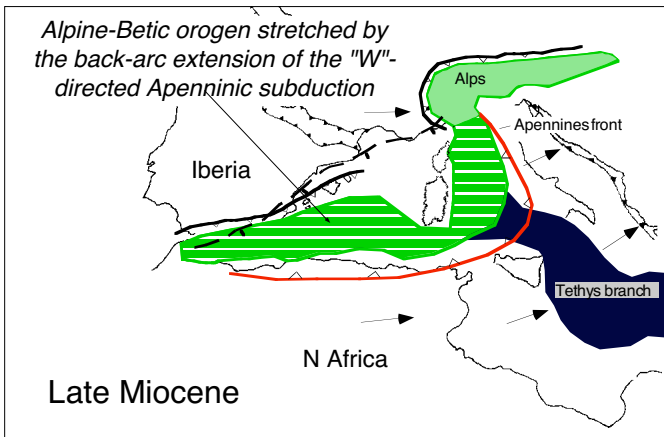
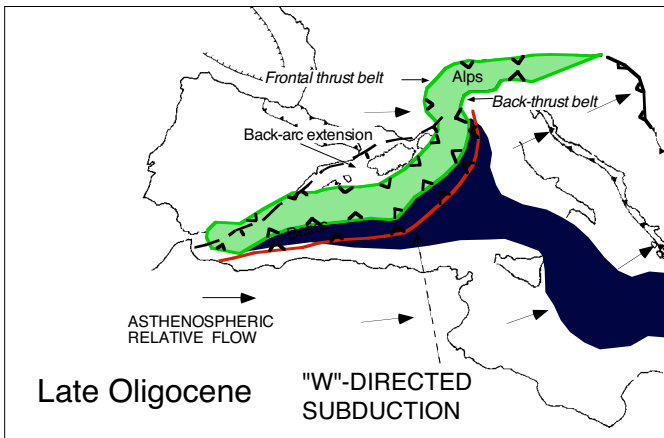
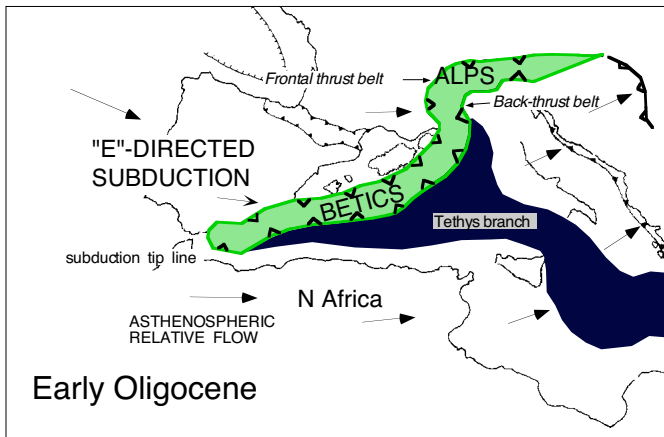
kinematics of the extensional and compressional belts of W-directed subduction zones



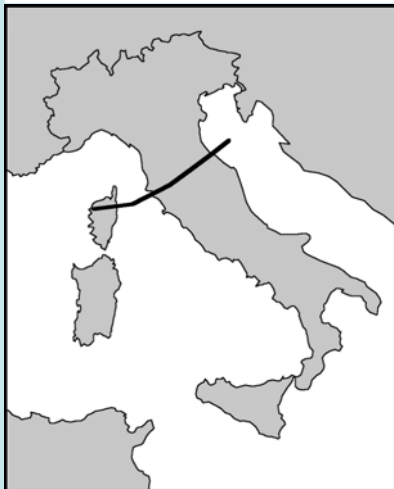
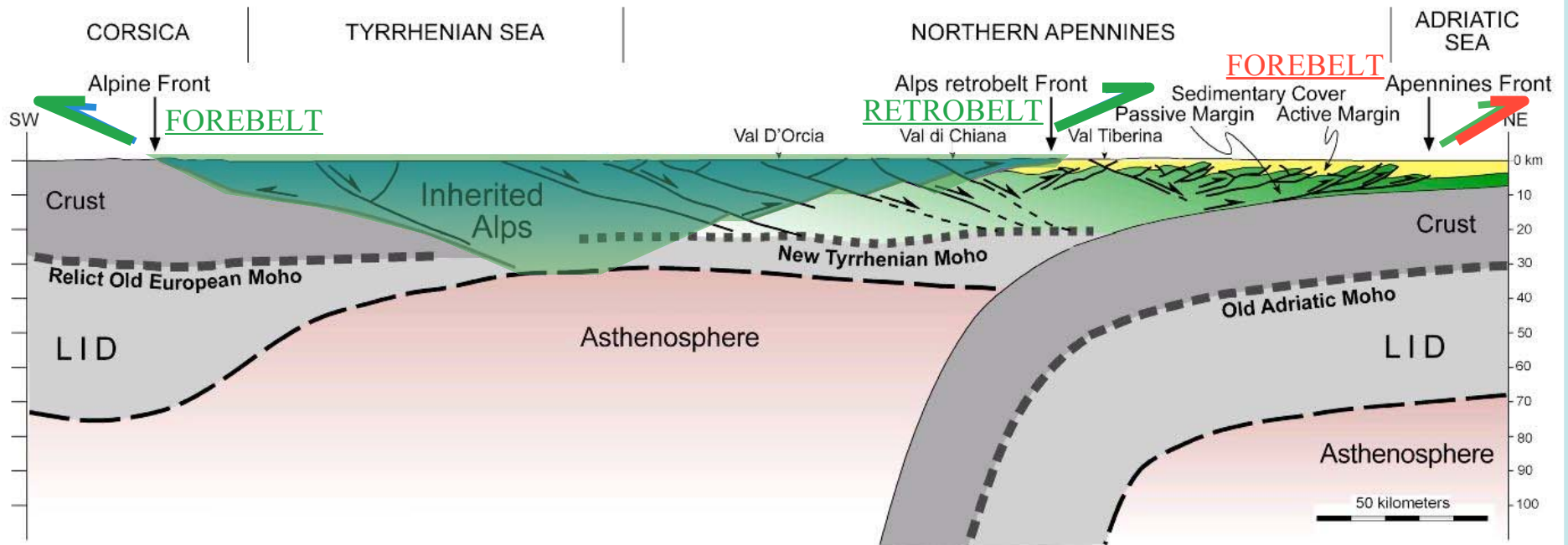


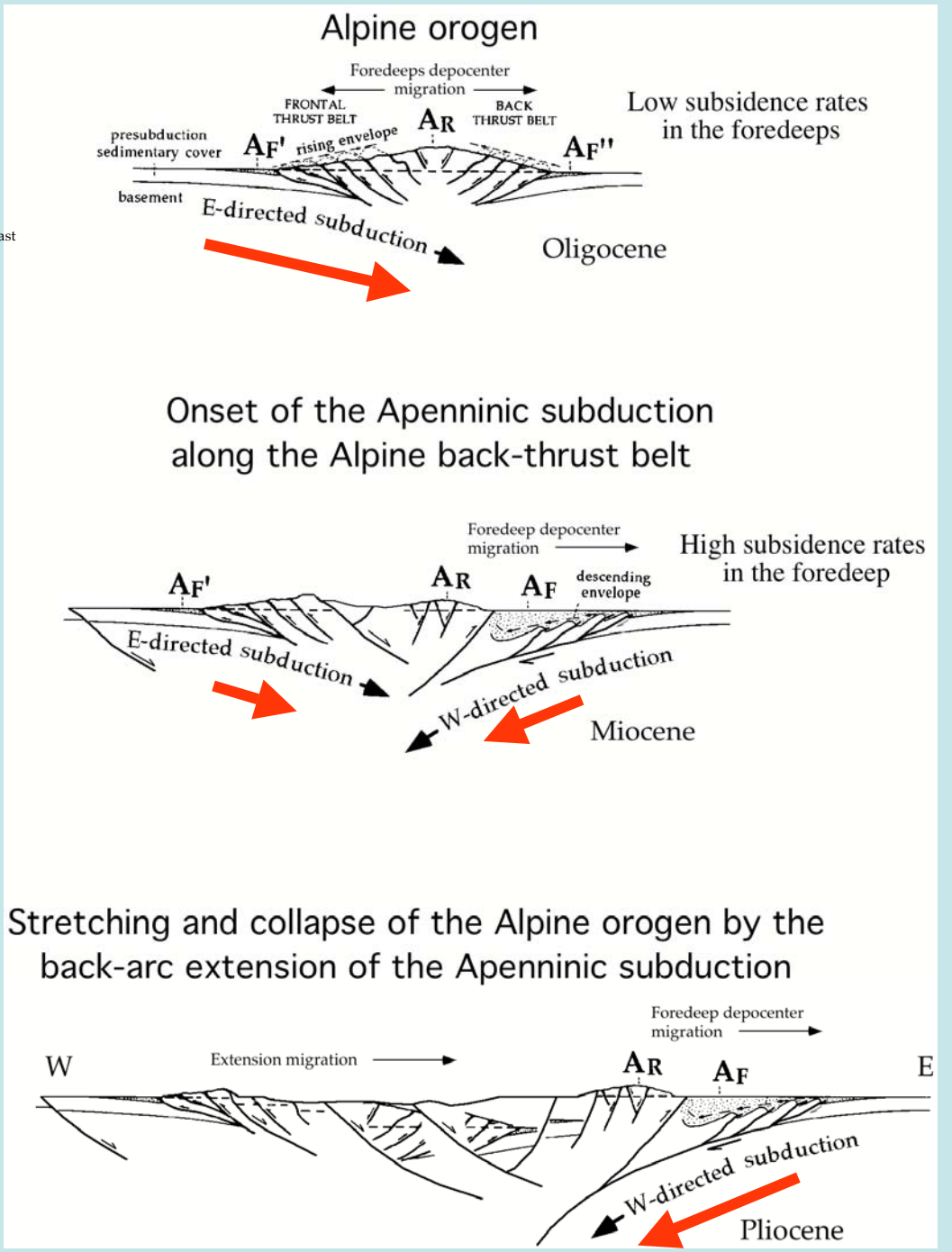
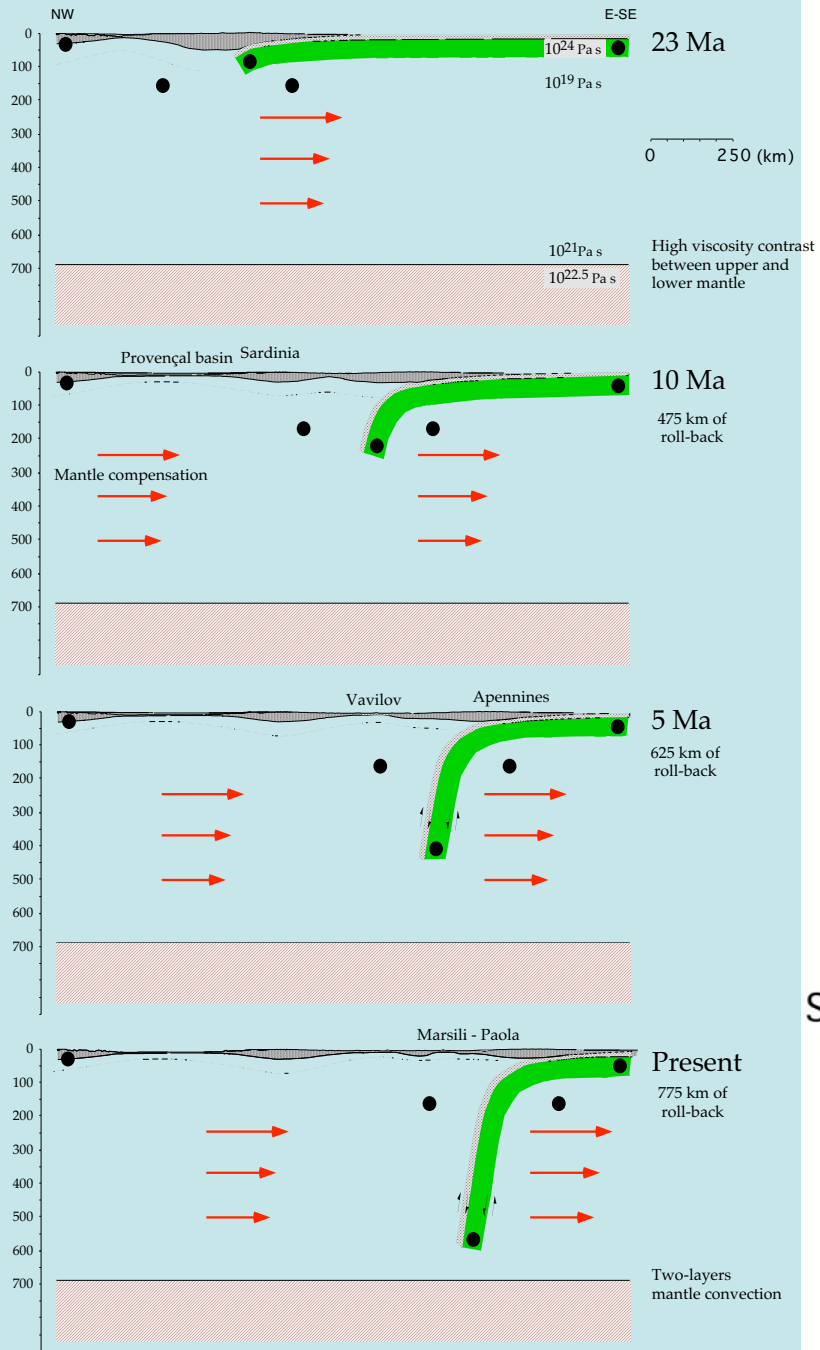


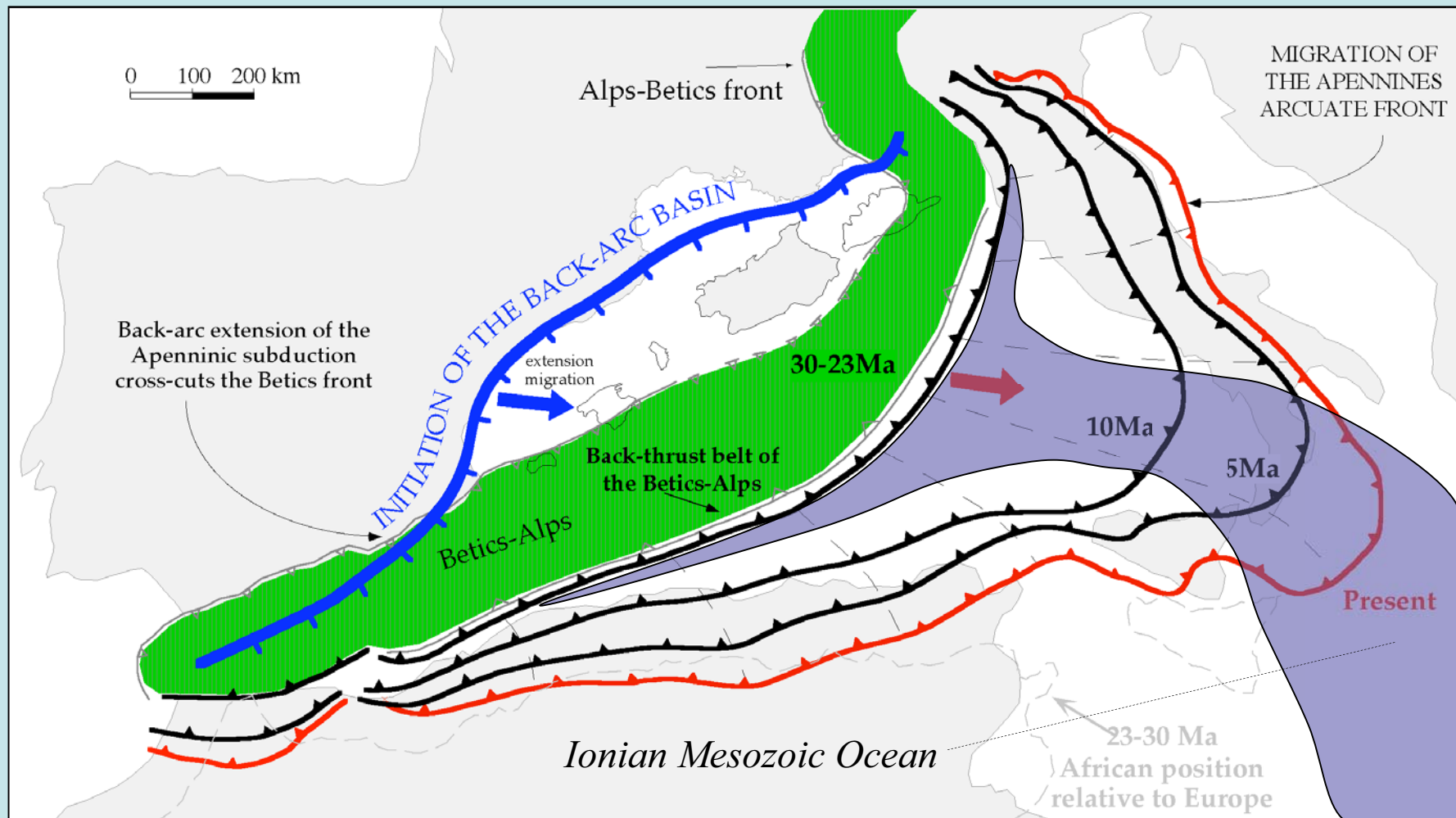




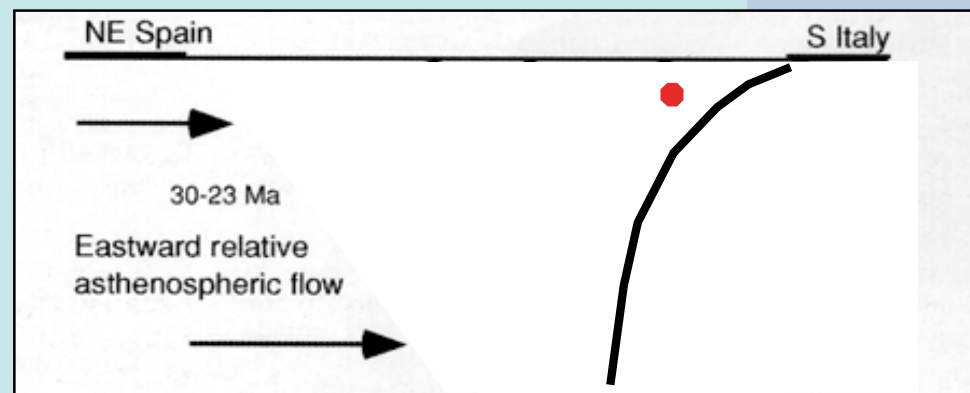
THE ALPS IN THE APENNINES







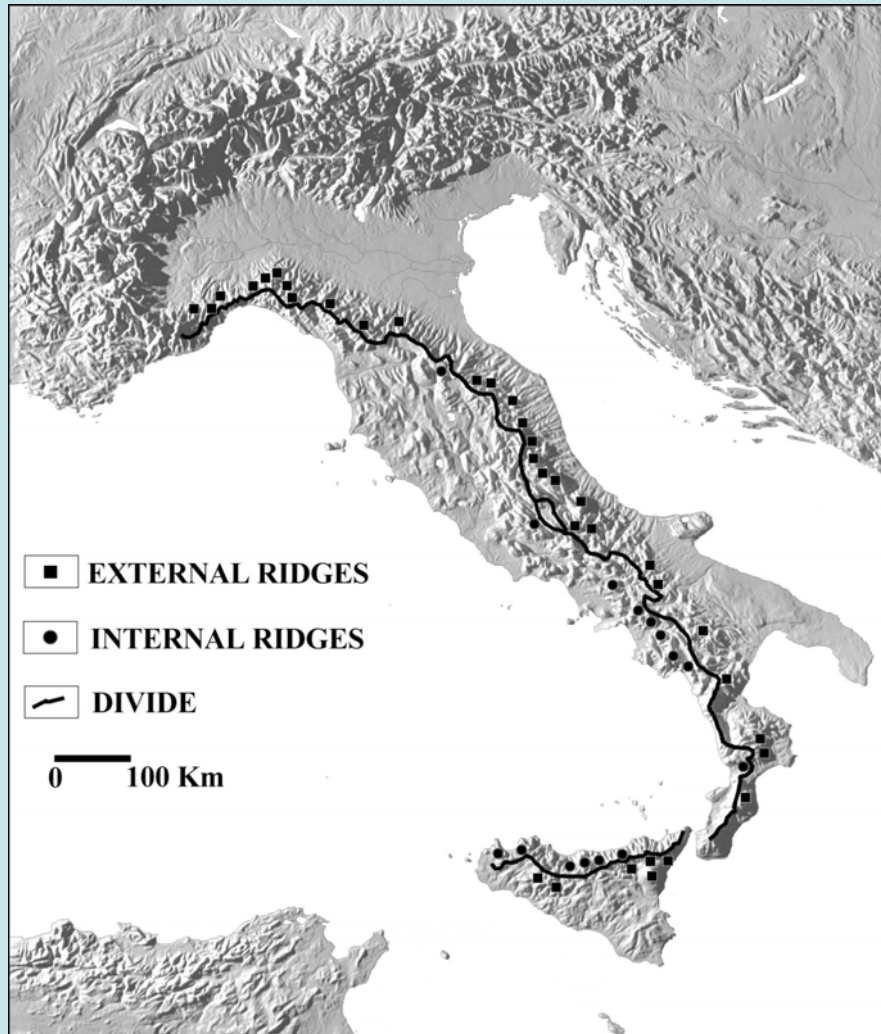
*E-ward slab retreat
5 times faster than N-S
convergence*



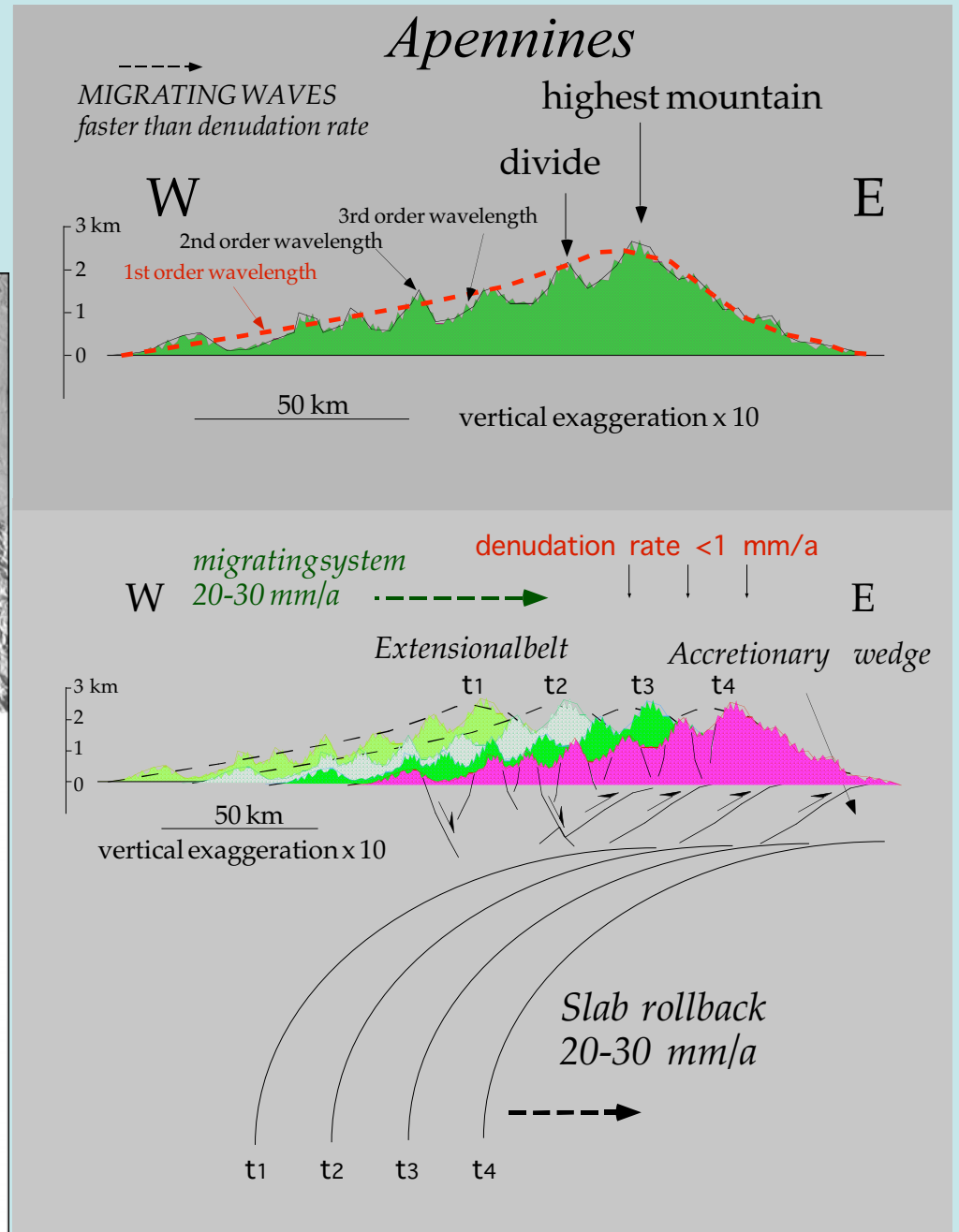
Water divide Vs Highest peaks



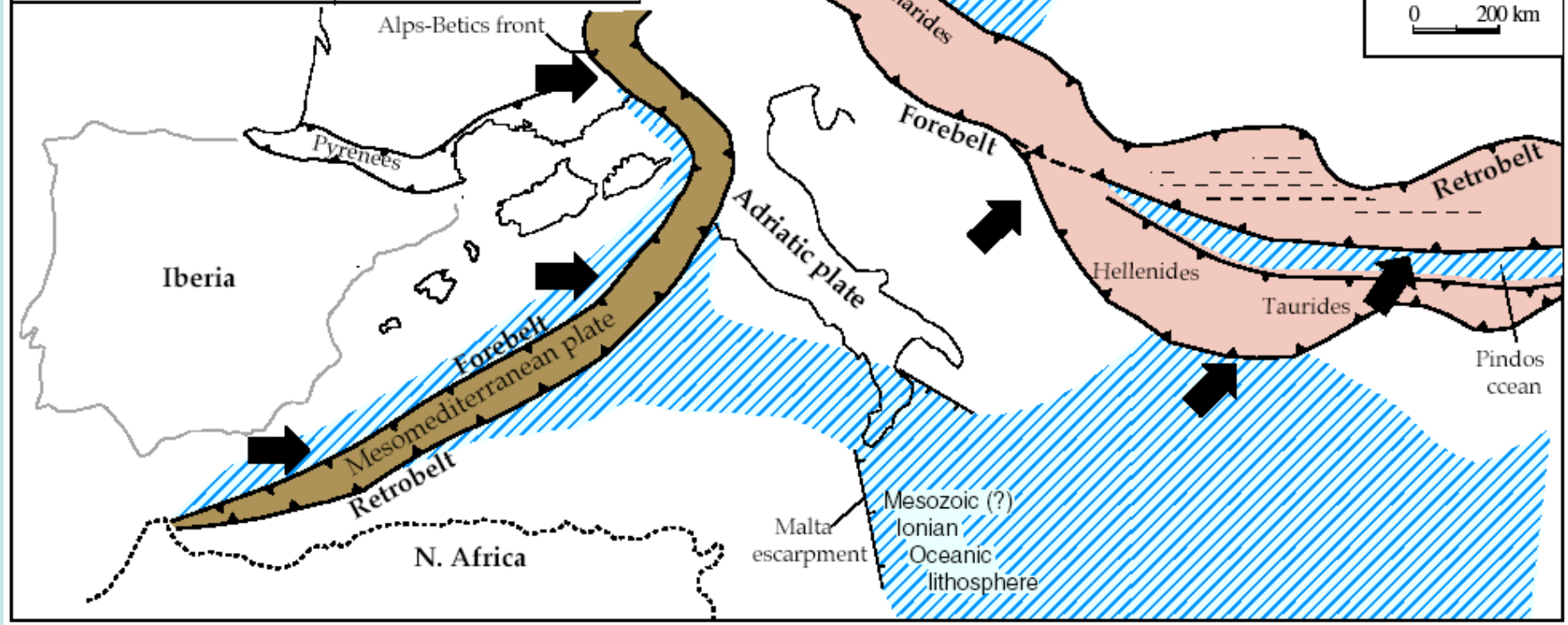
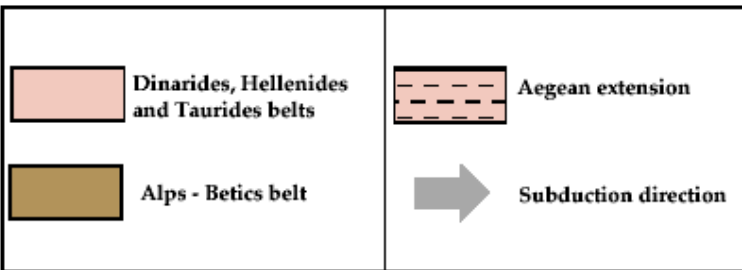
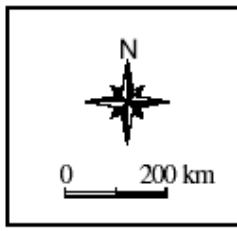
Water divide Vs Highest peaks

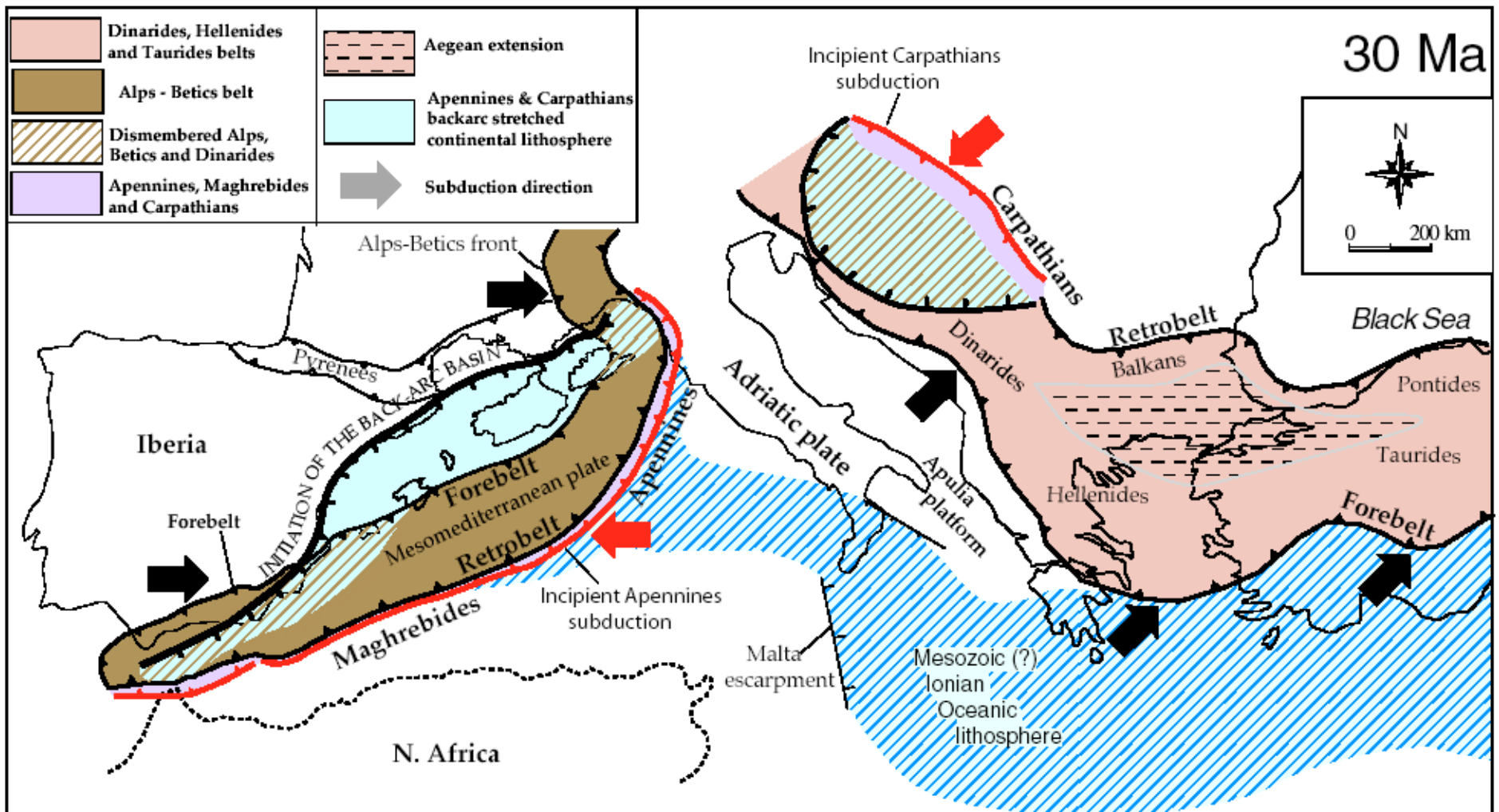


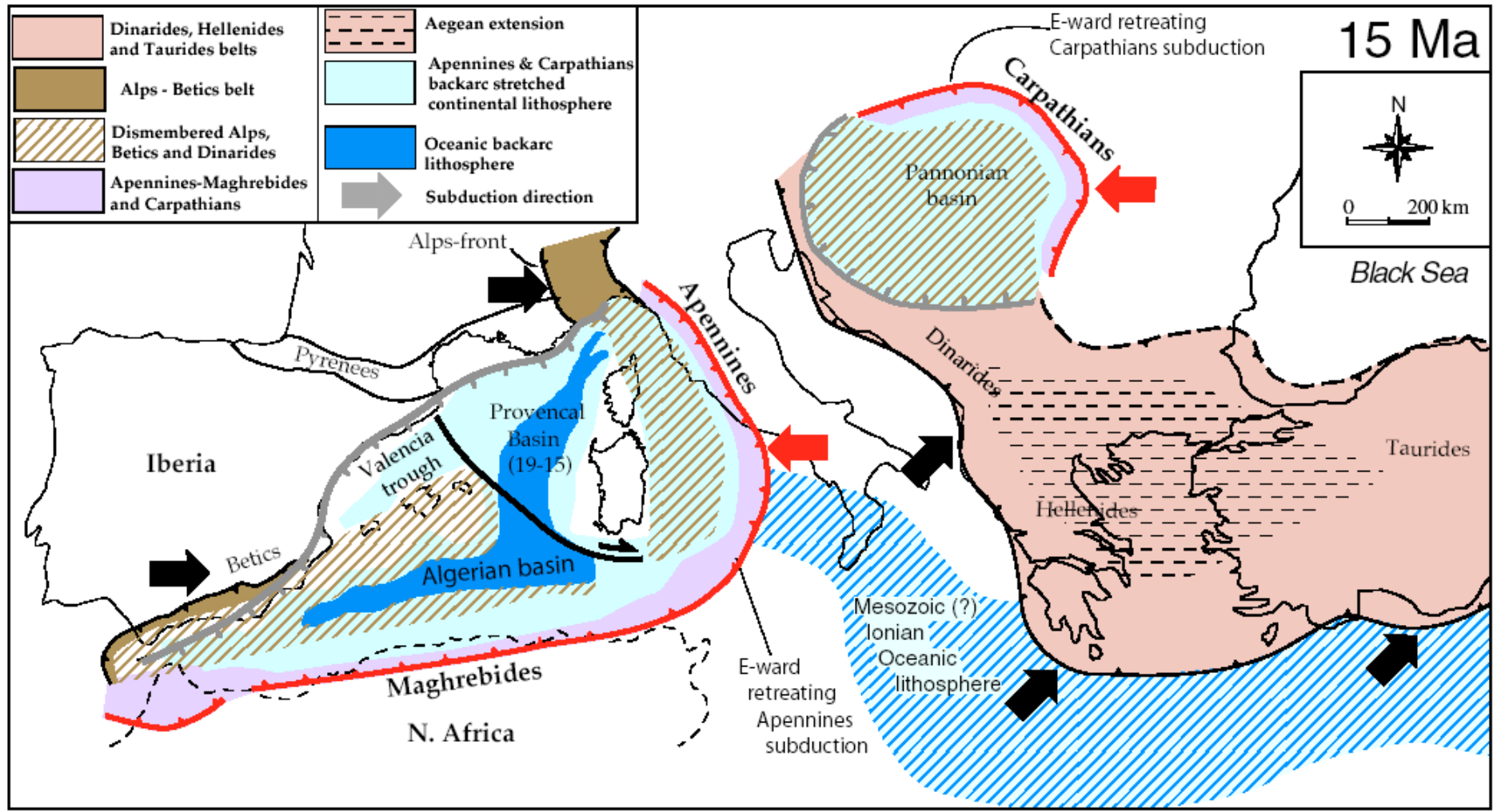
Salustri et al., 2002

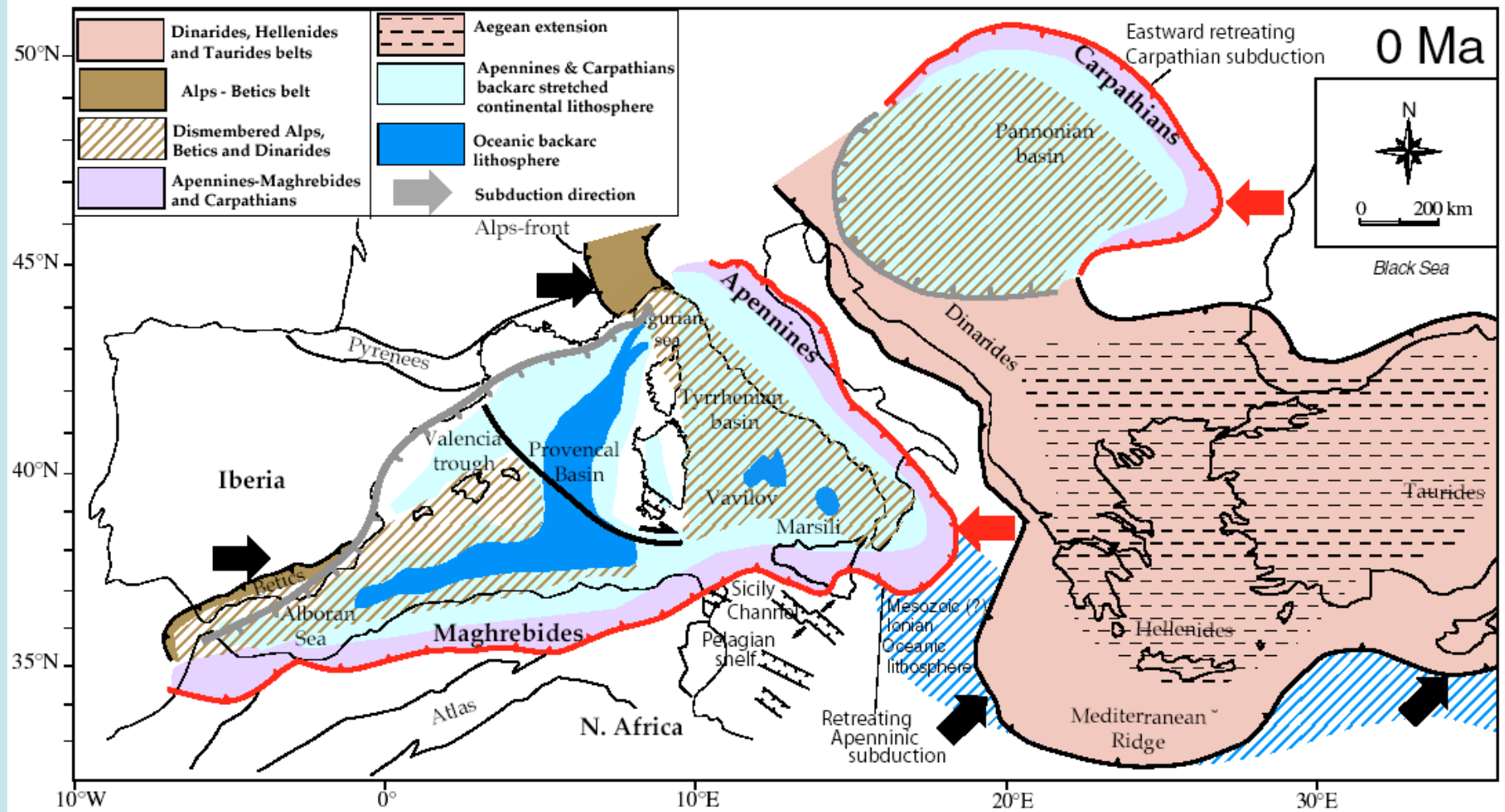


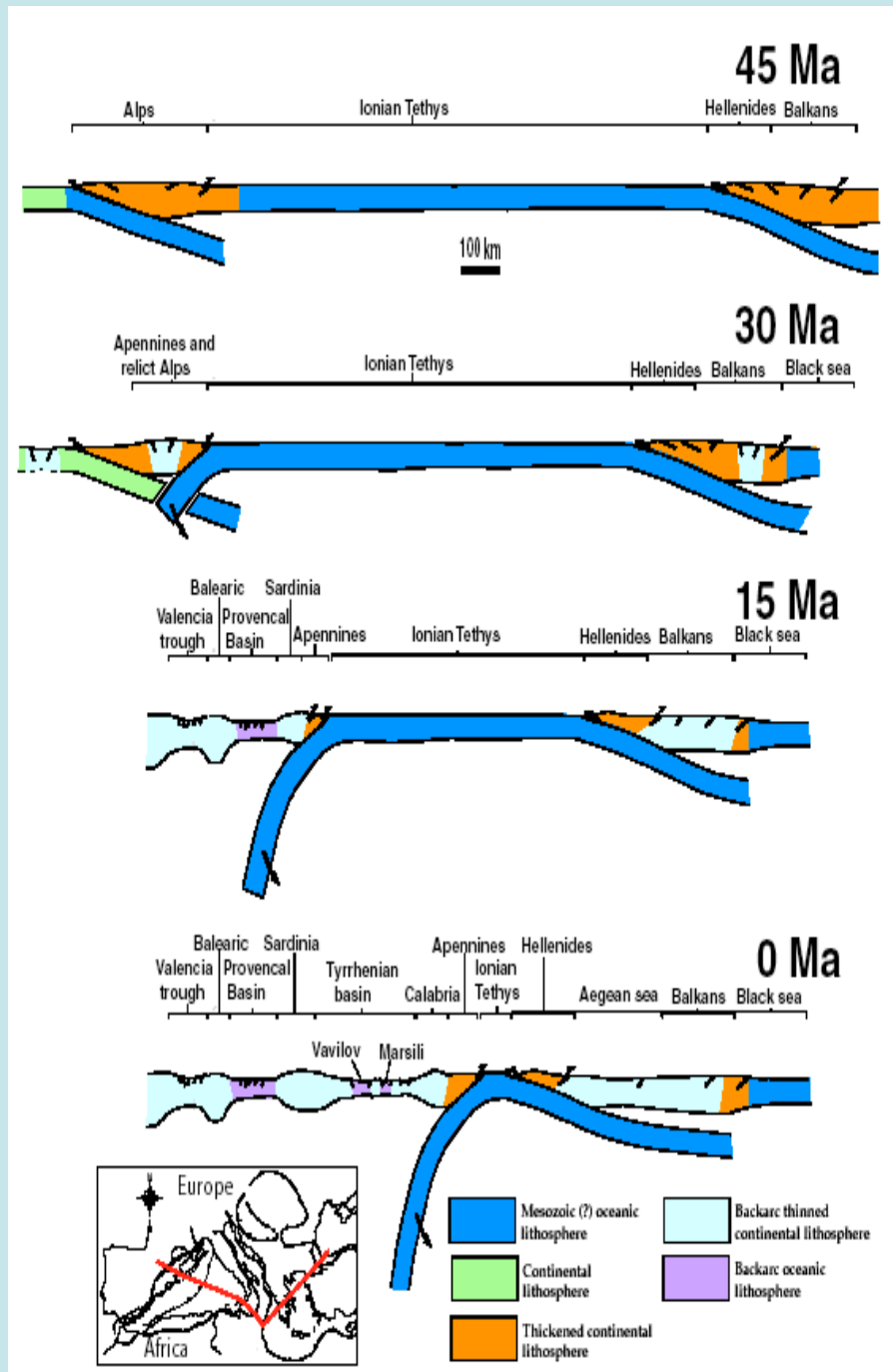
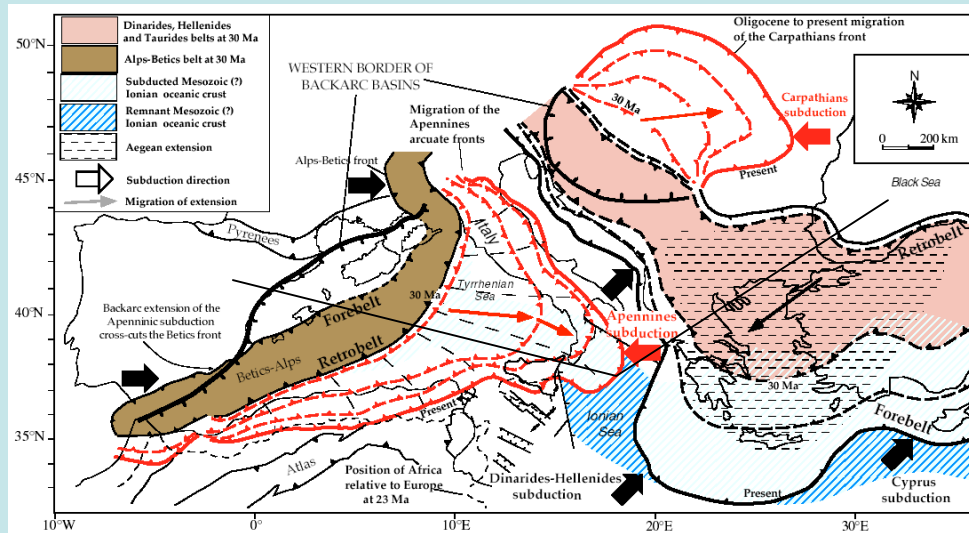
45 Ma

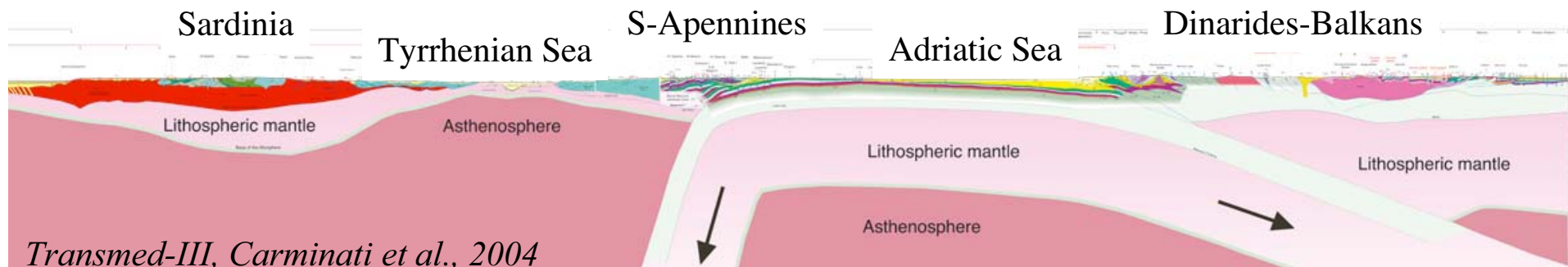
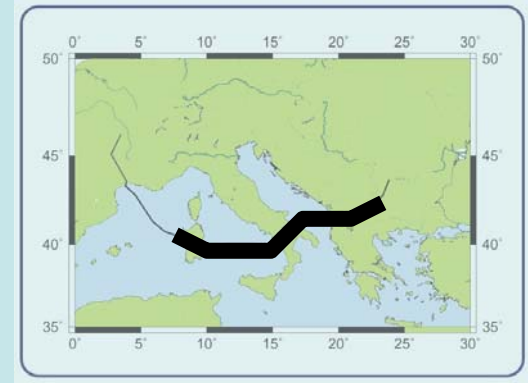






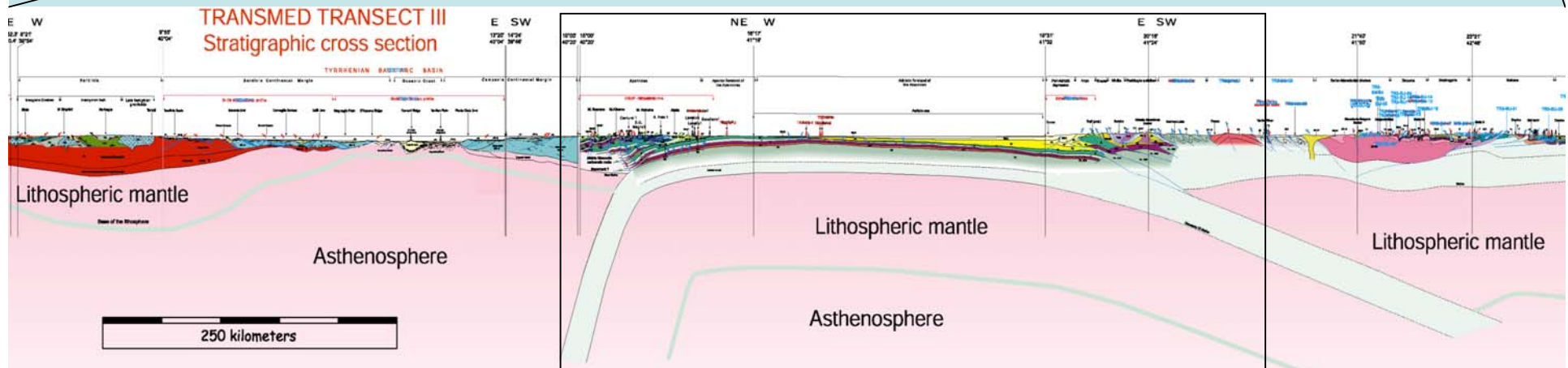
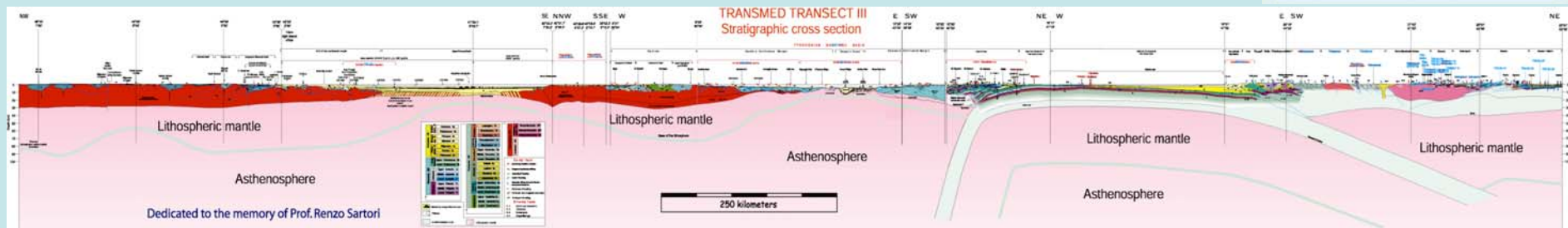
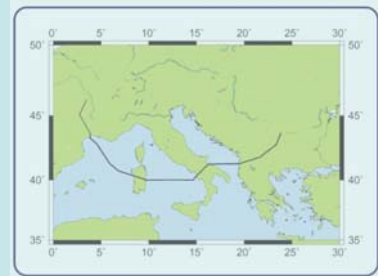






Transmed-III, Carminati et al., 2004

250 km



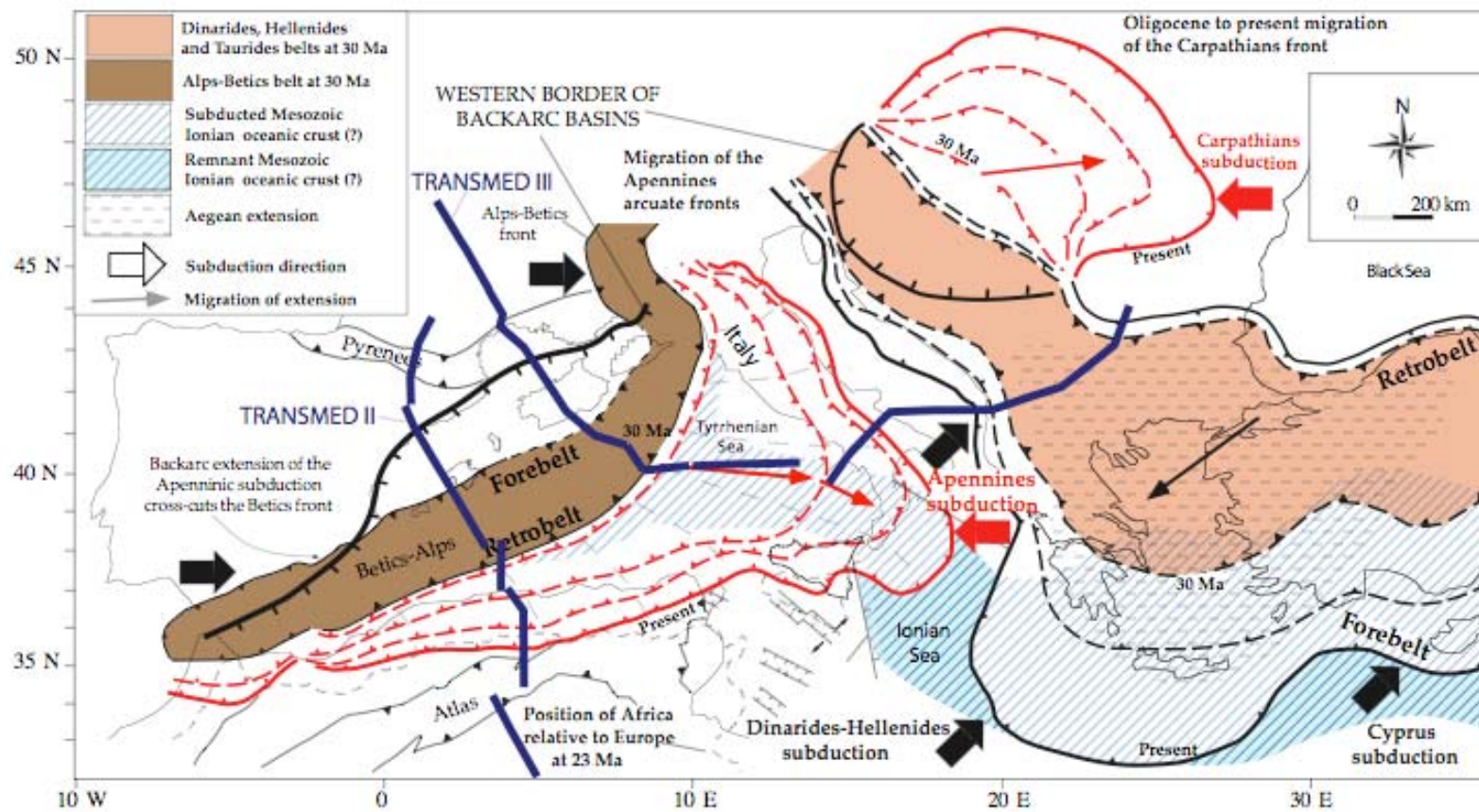
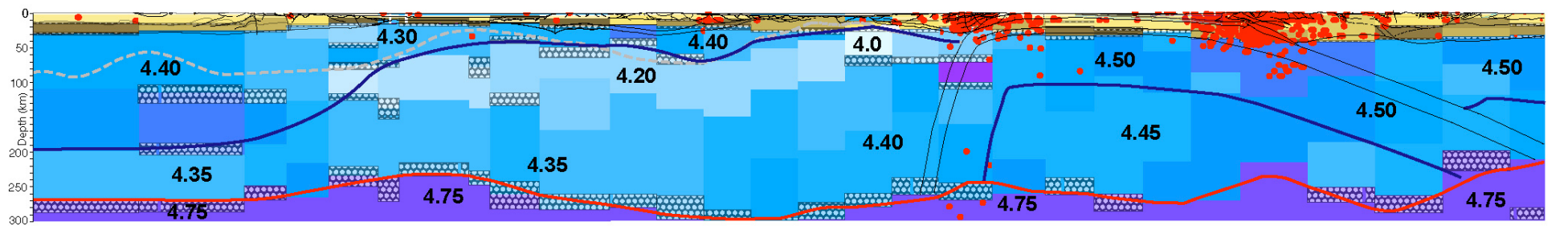
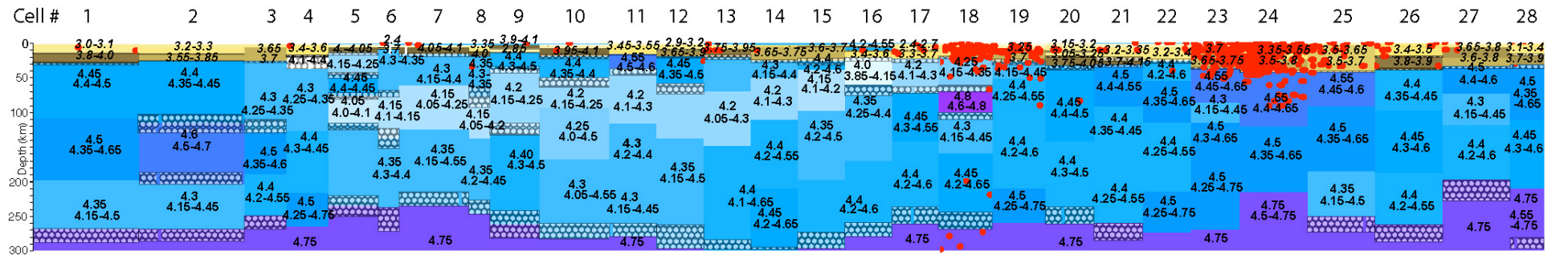
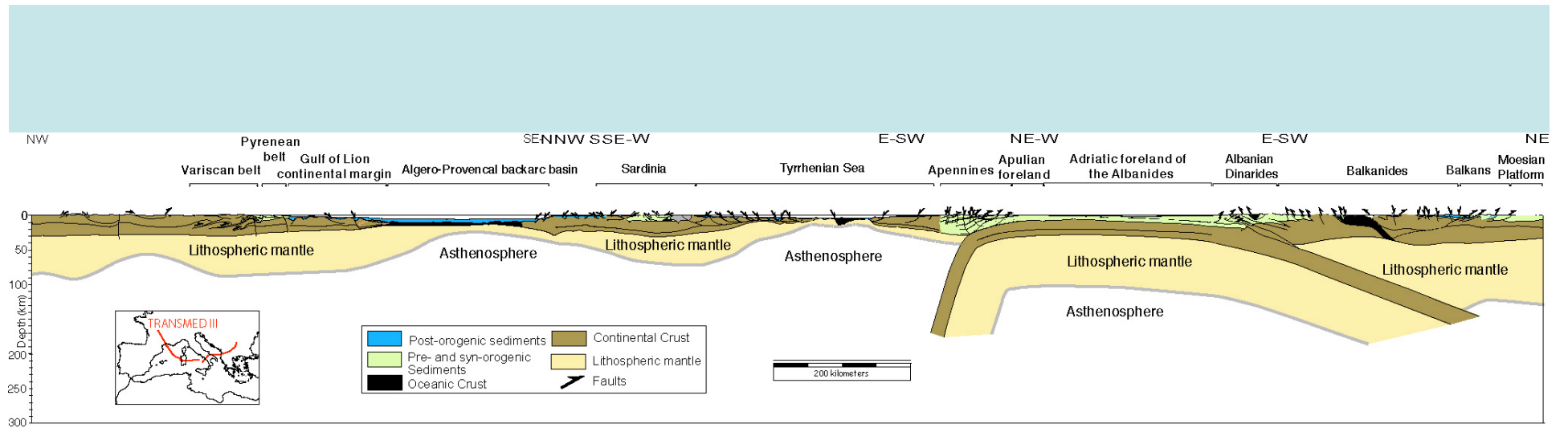
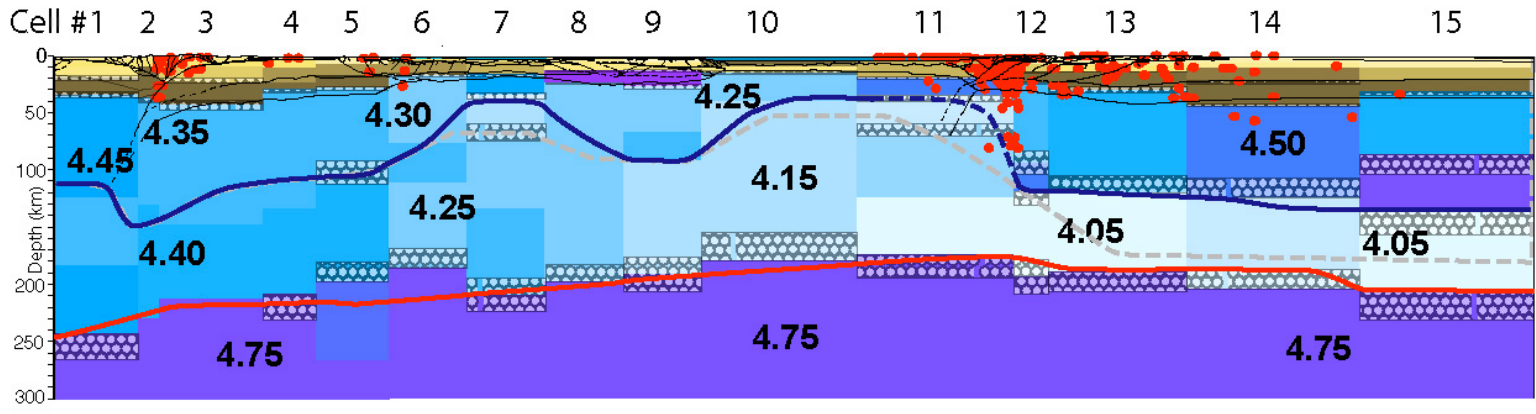
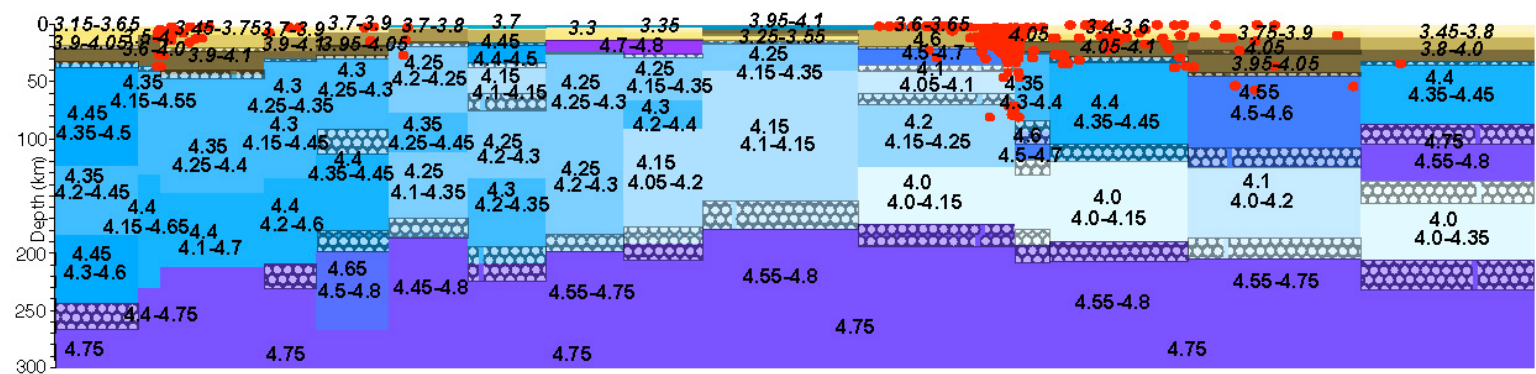
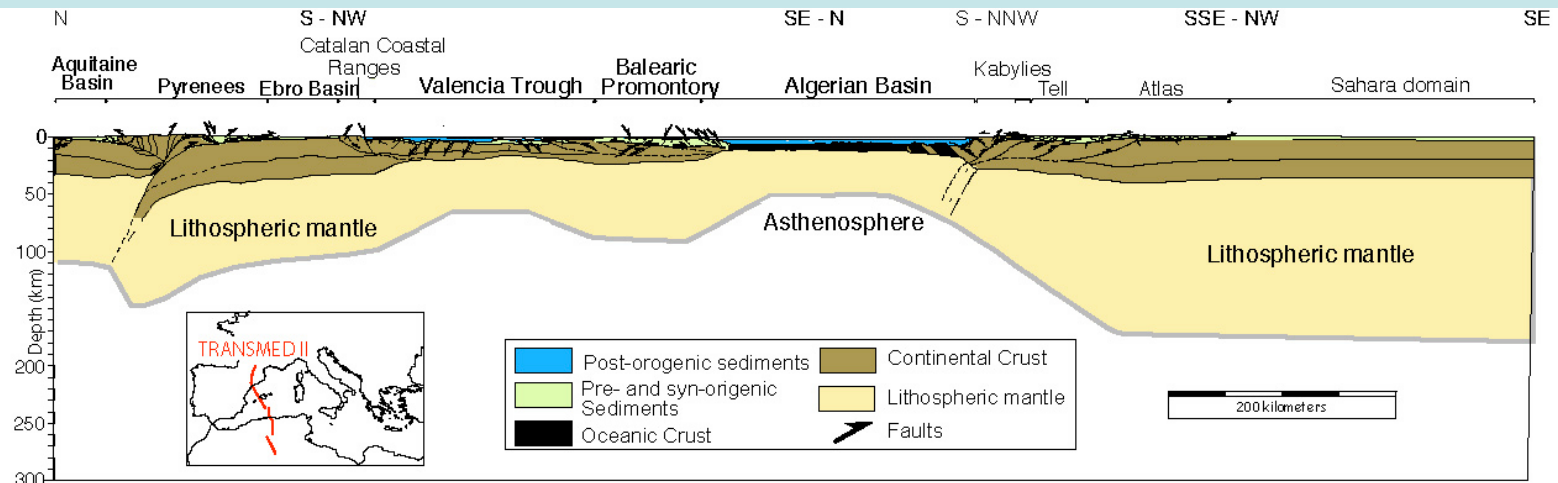
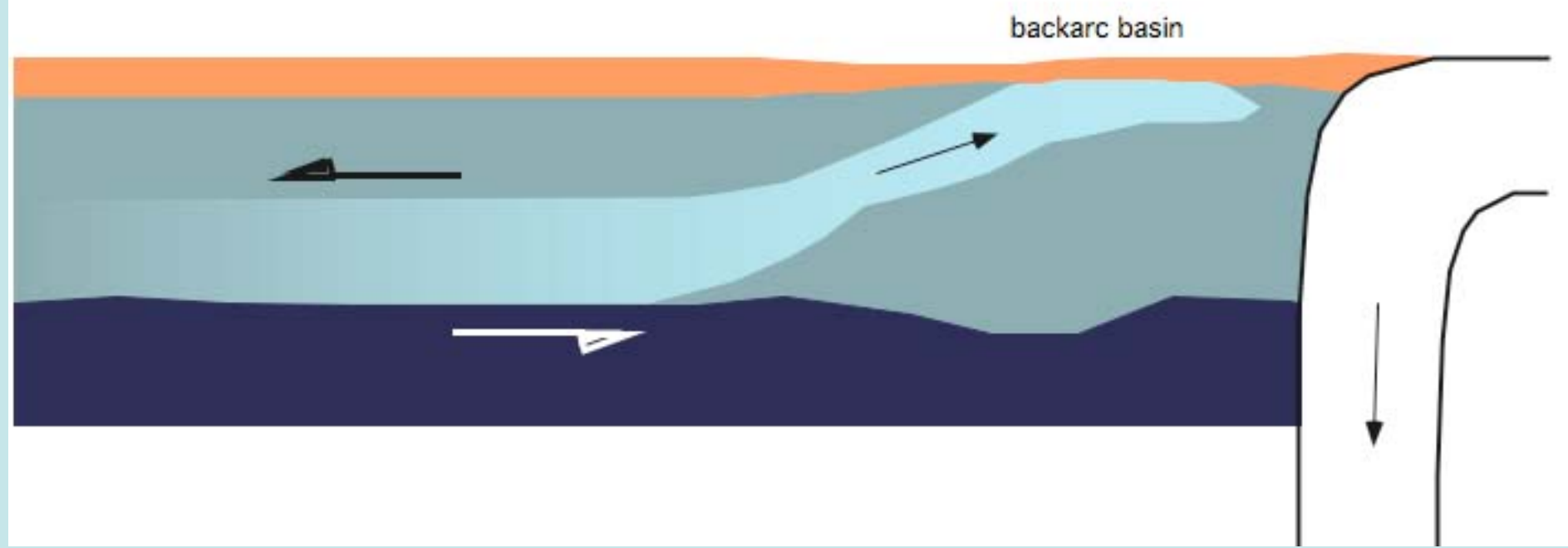
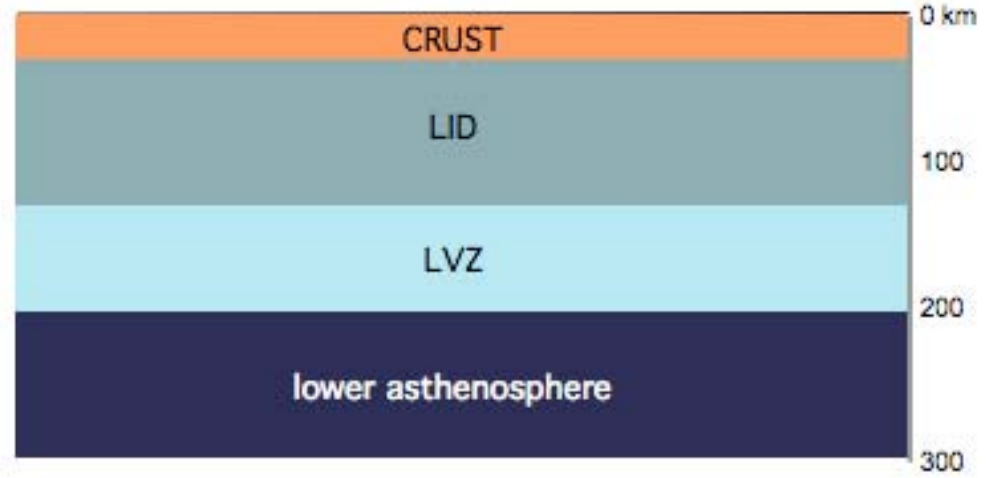


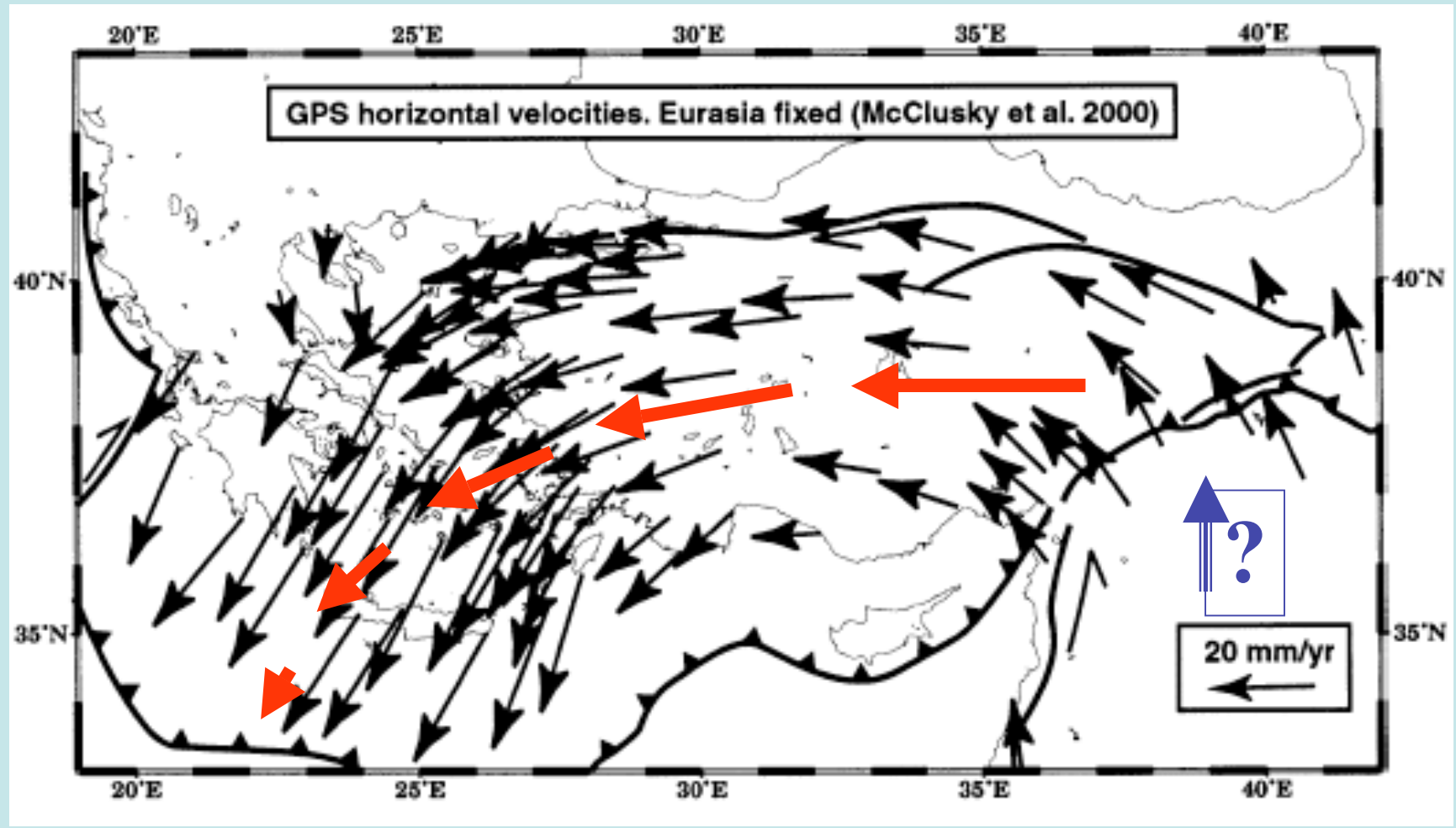
Fig.1



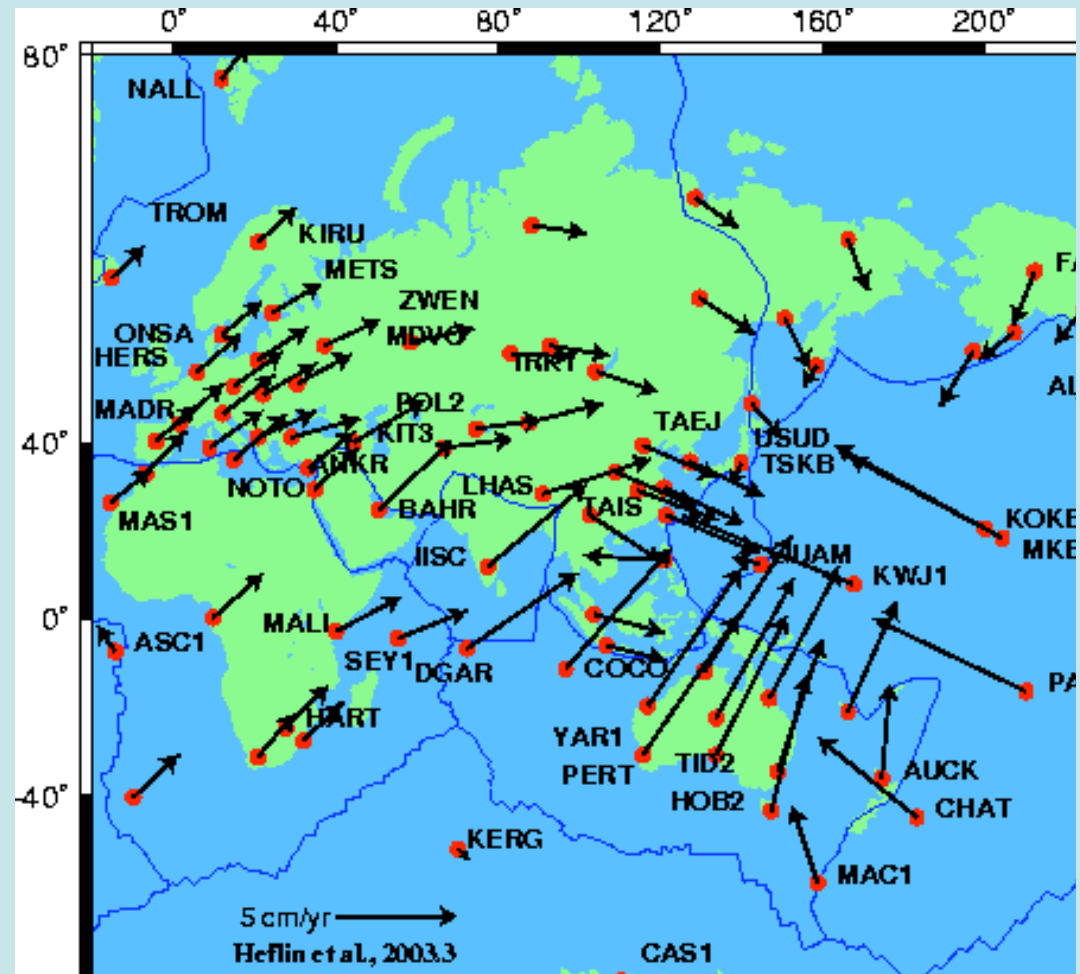


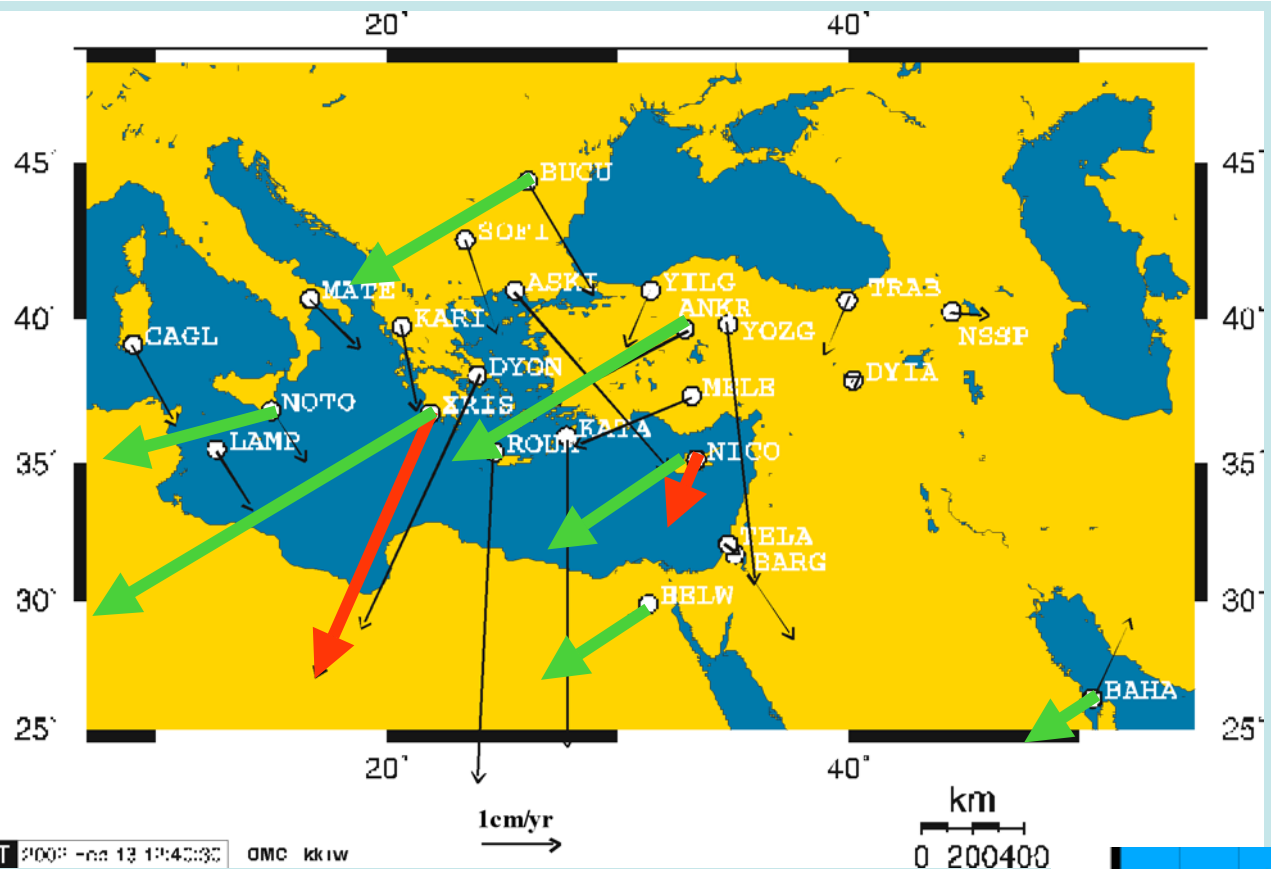
Panza et al. 2006





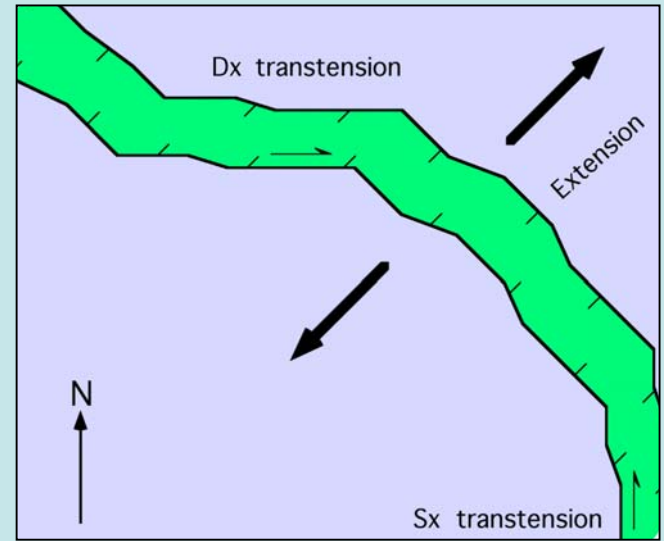
*E-Med subductions relate
Greece & Turkey to Africa*

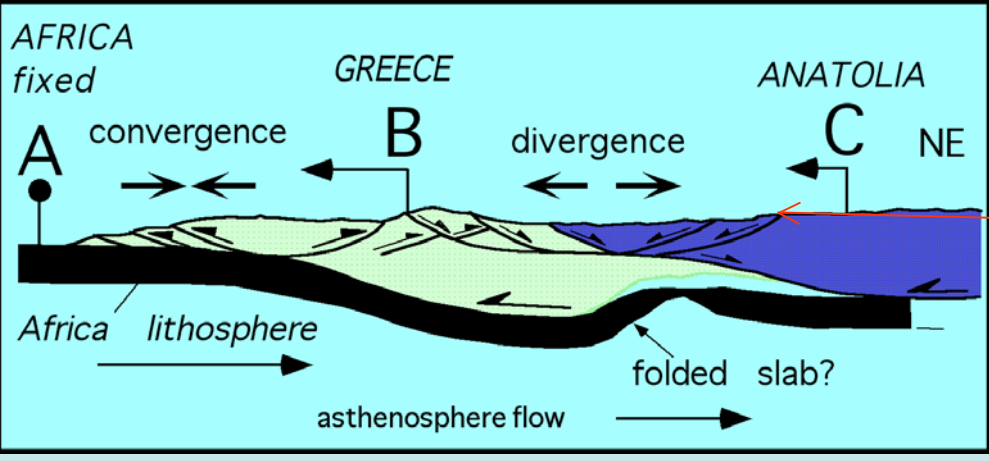
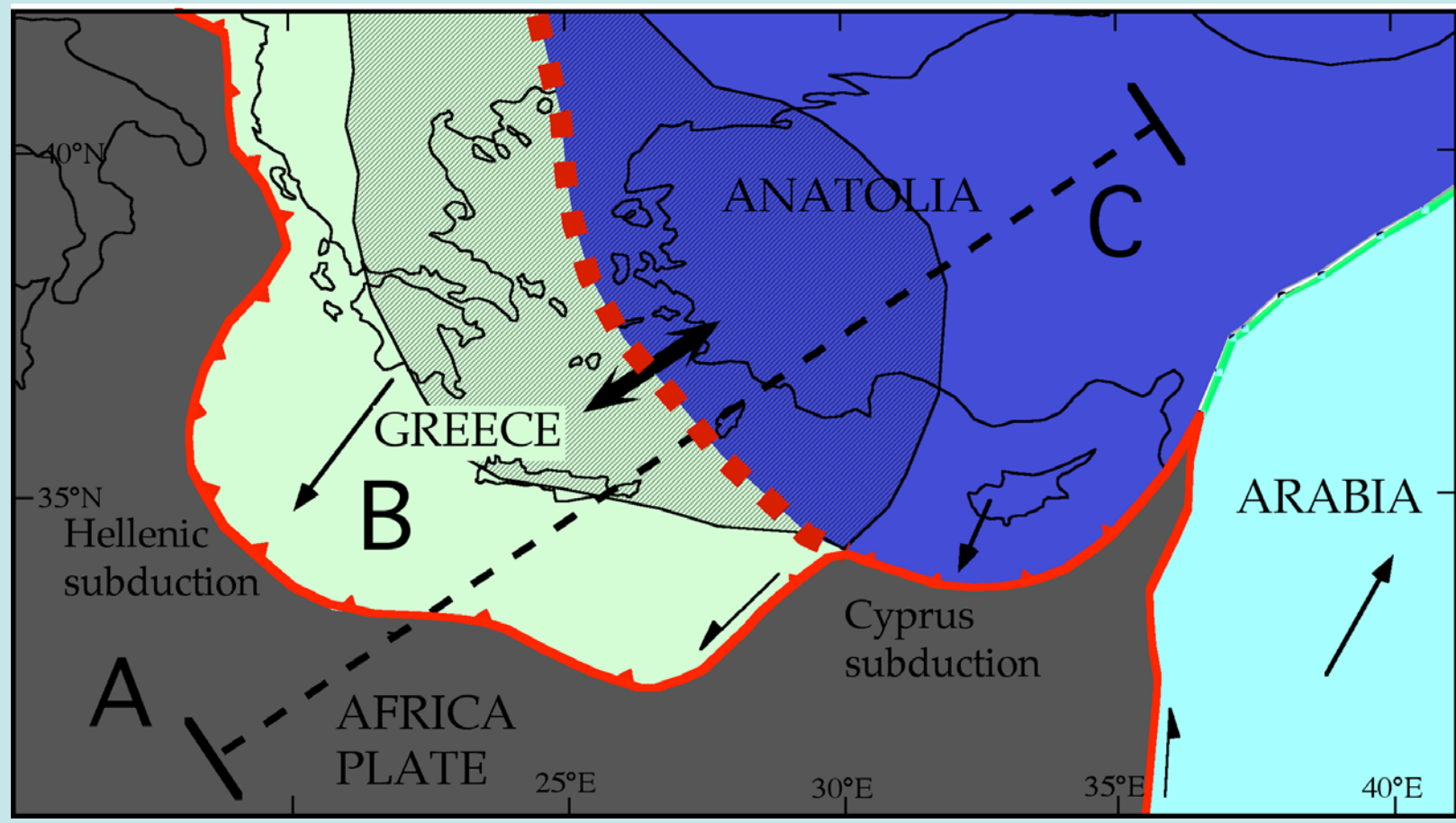


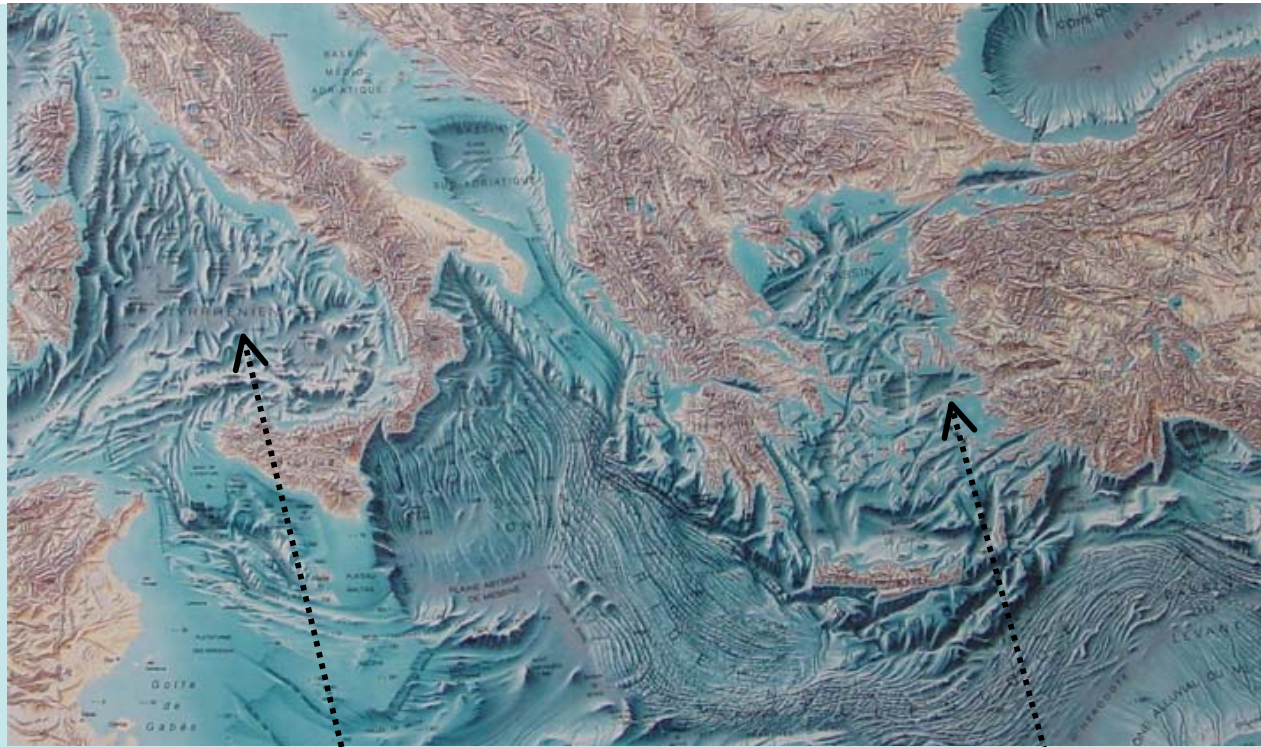


 *Africa fixed*
 *Hawaii fixed*

GMT 2008 -aa 13 12:43:33 OMC kkiw





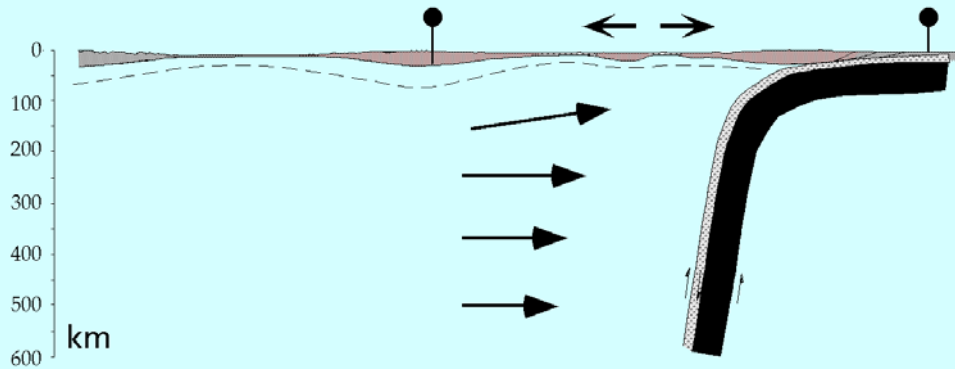


Thin lithosphere

Thick lithosphere

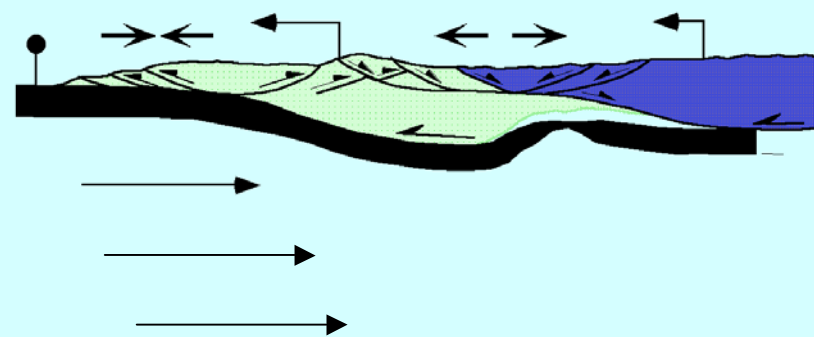
1 plate

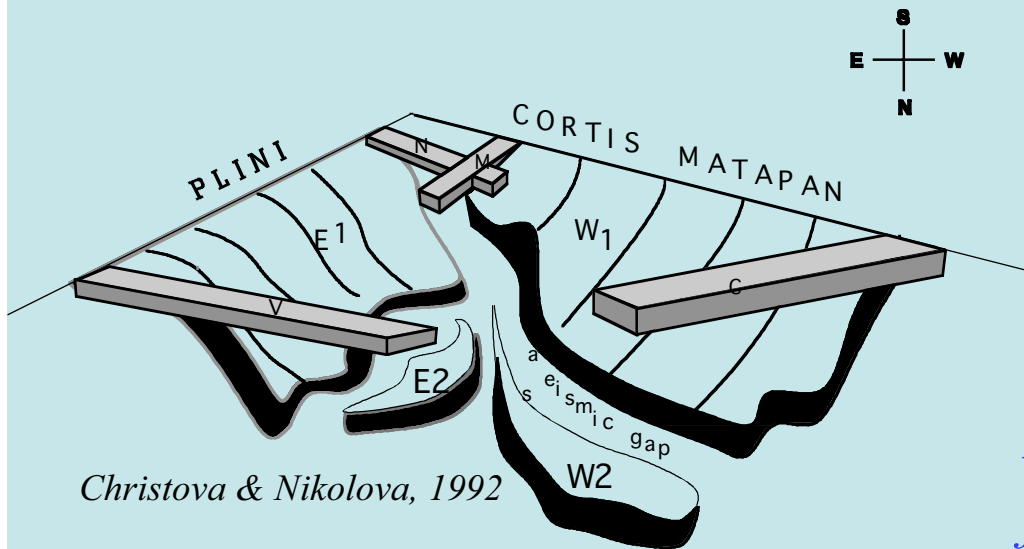
Asthenospheric upwelling replacing slab retreat



3 plates

Hangingwall extension in spite of lithospheric thickening

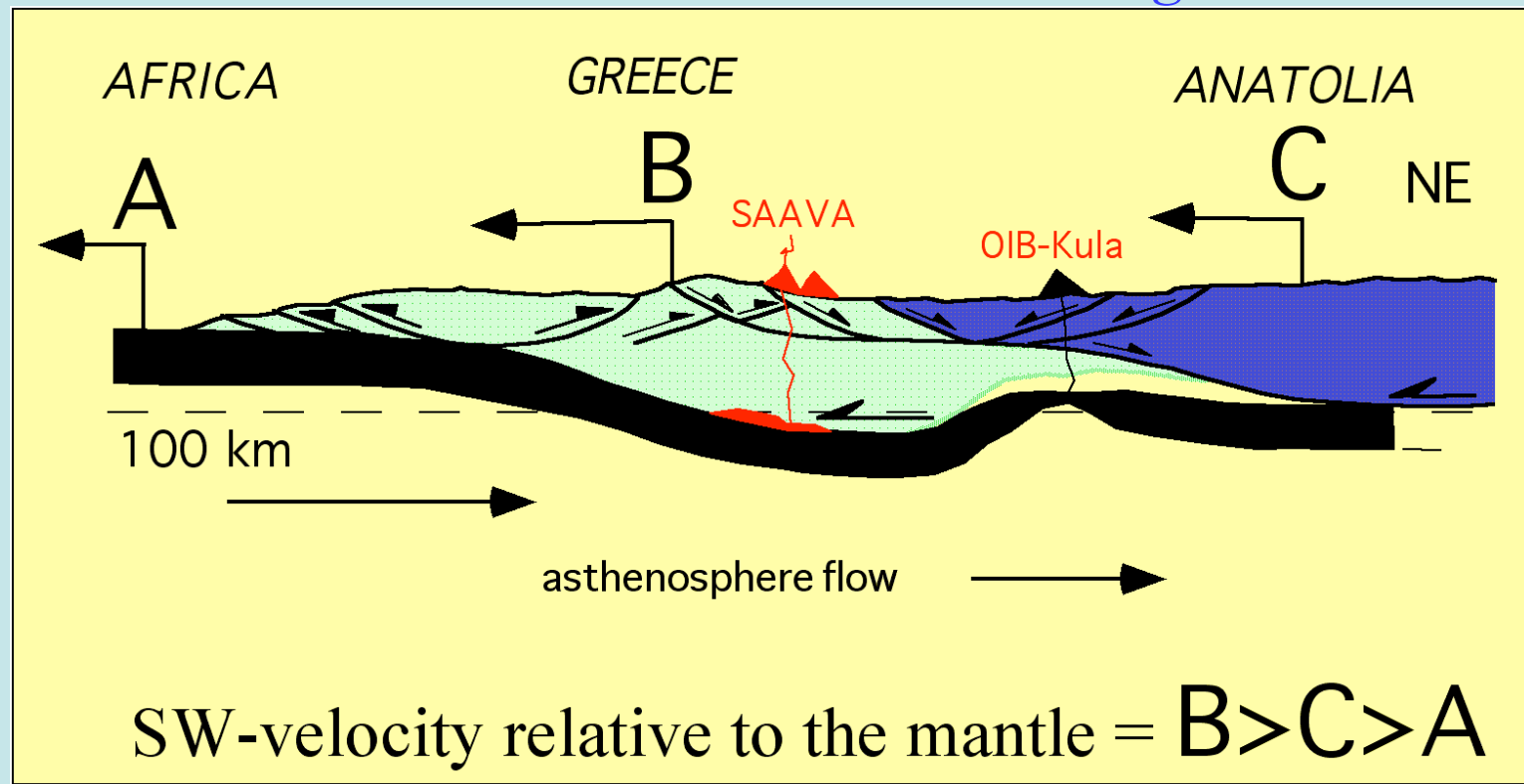




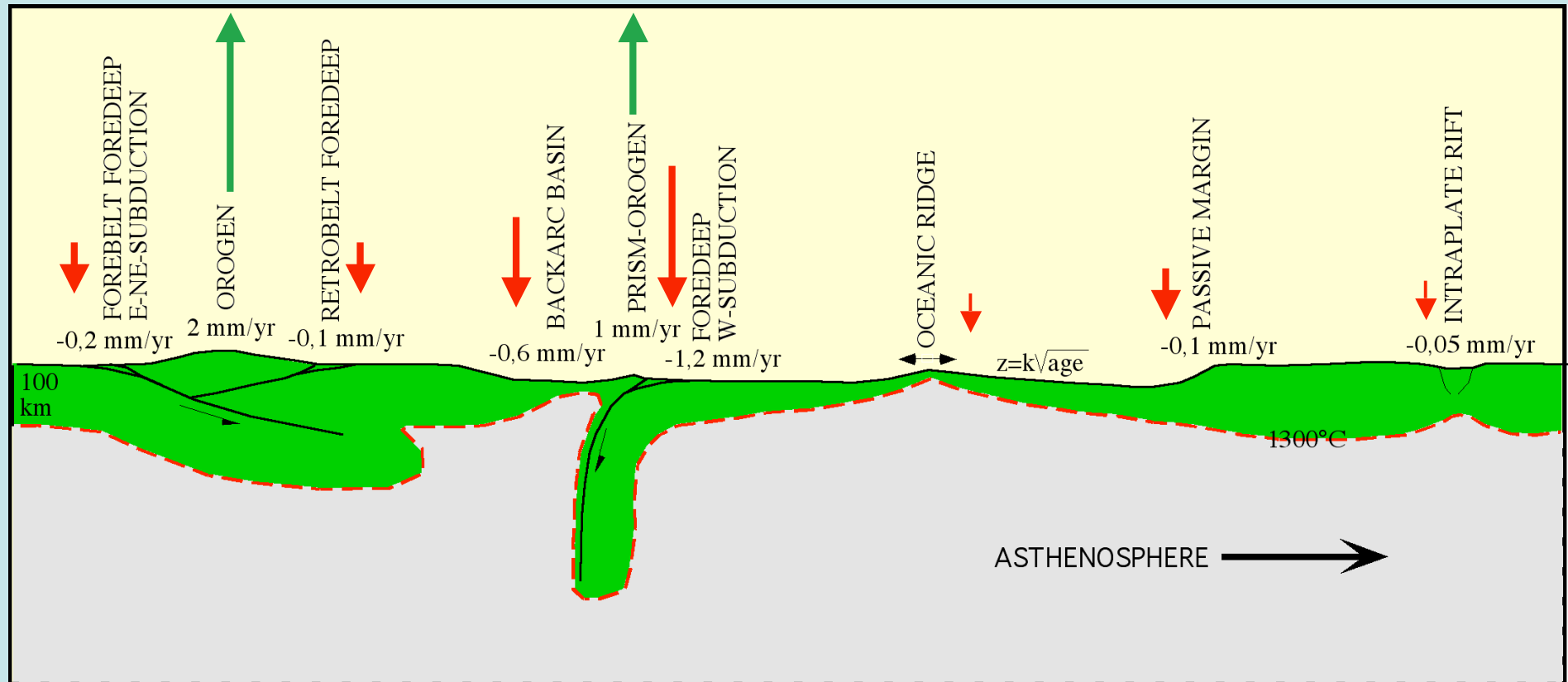
Christova & Nikolova, 1992

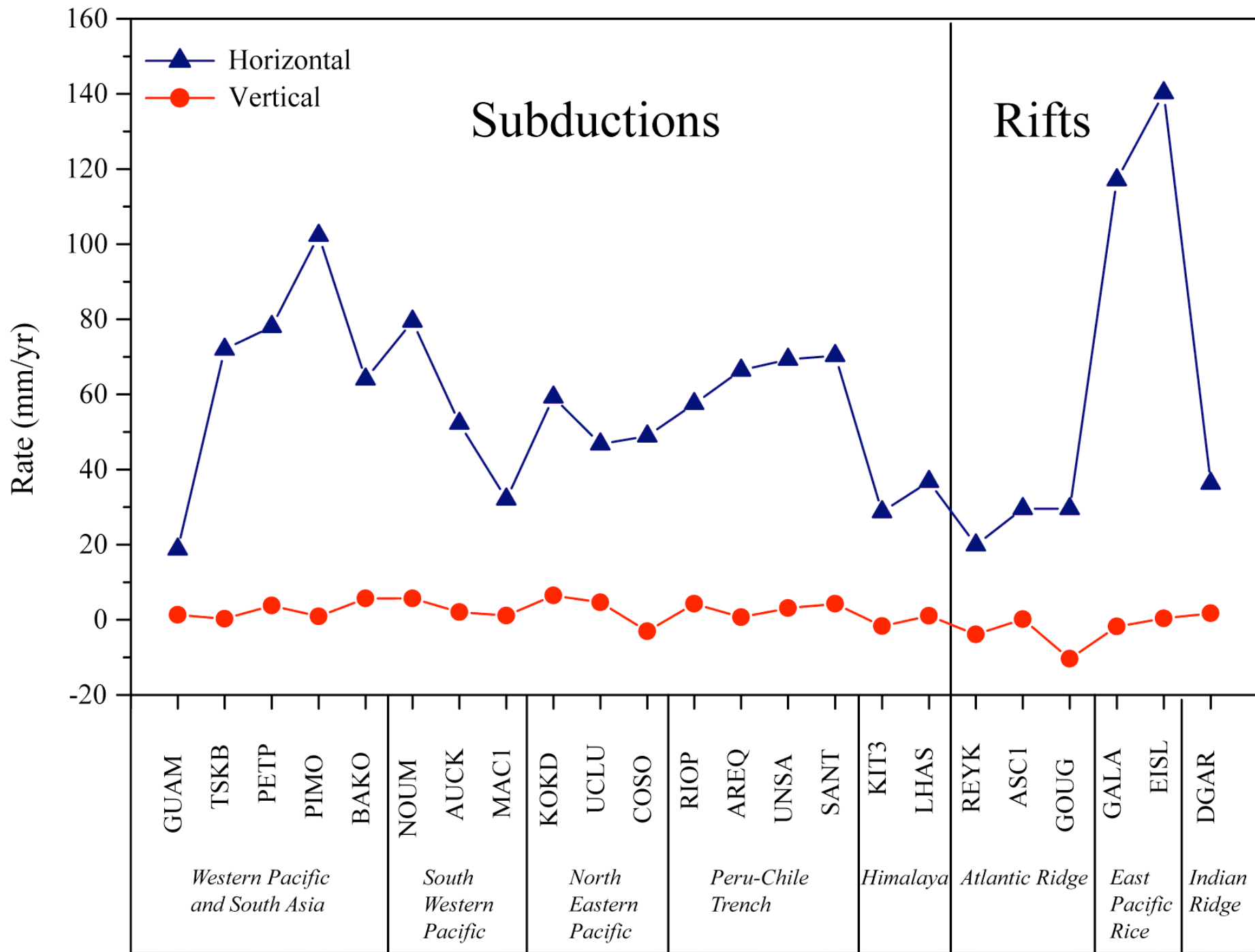
**FASTER CONVERGENCE =>
LARGER MAGMATISM**

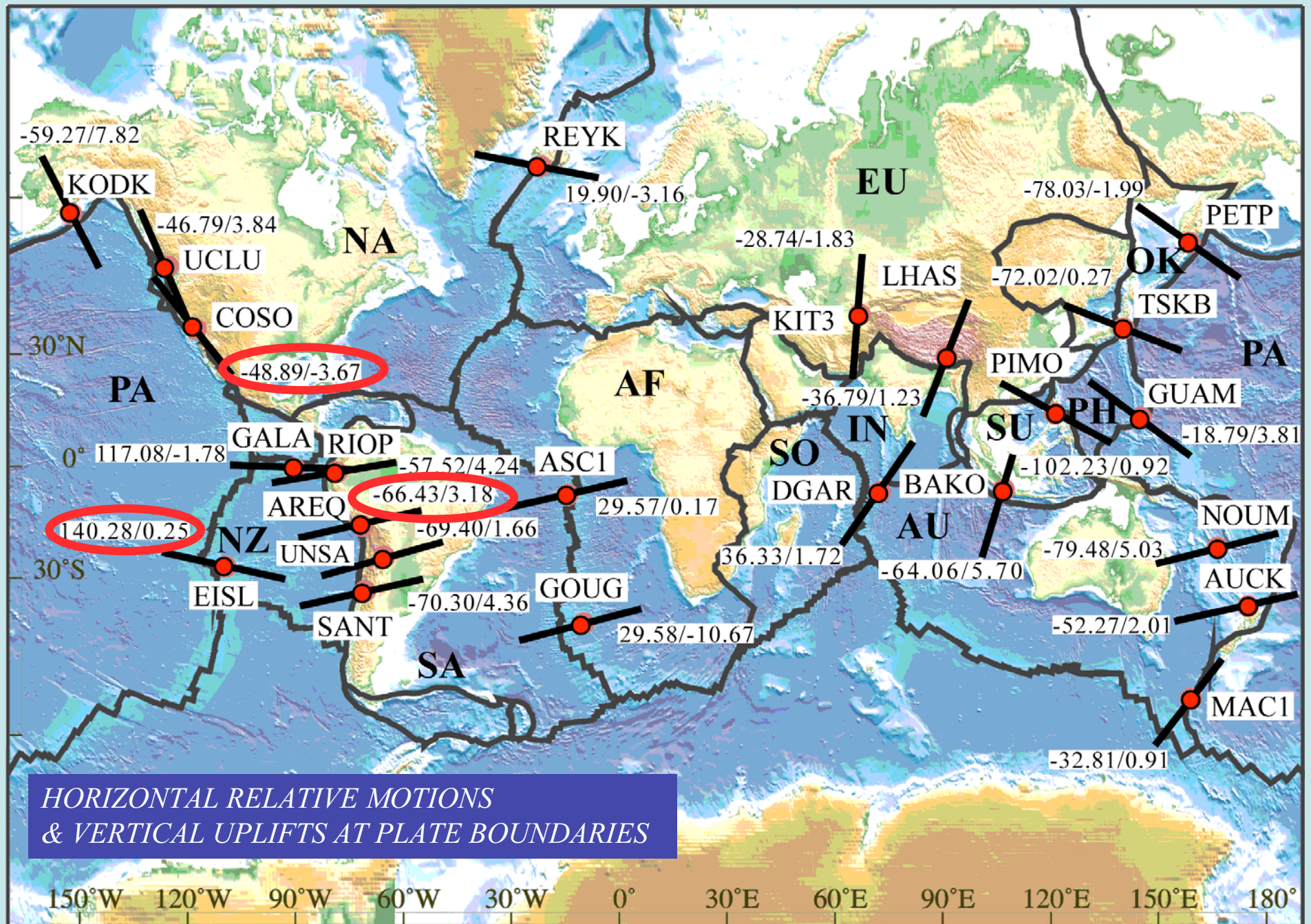
*A working hypothesis
for the SAAVA magmatism:
Shear heating ?*

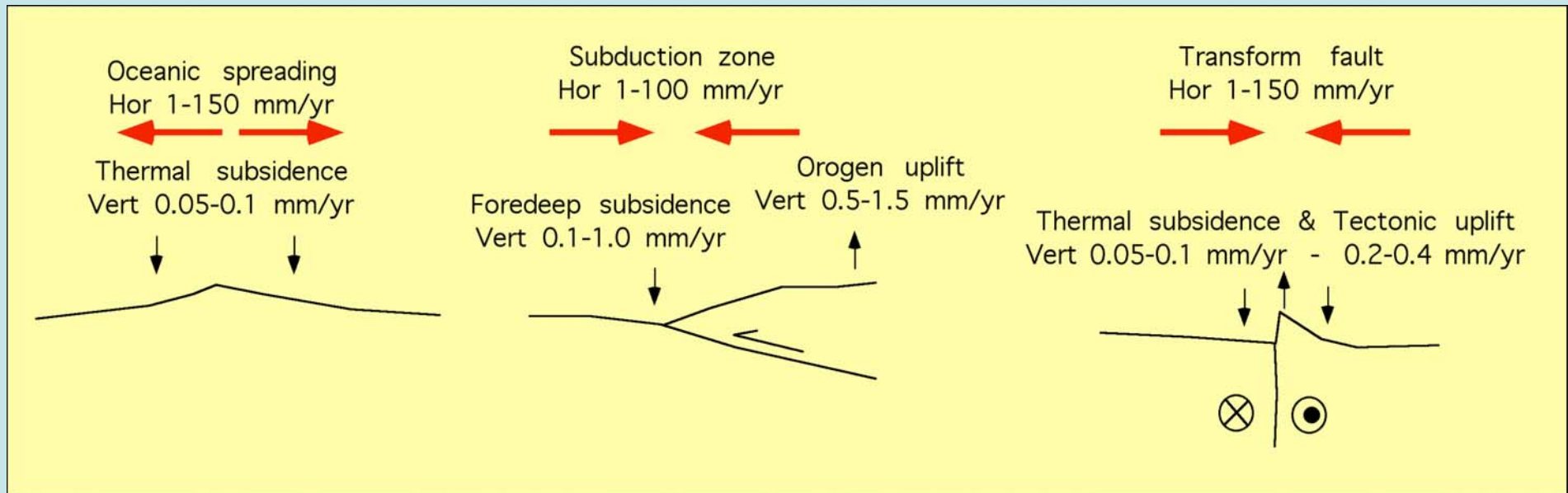


VERTICAL MOVEMENTS - PLATE TECTONICS RELATED









- *HORIZONTAL MOTIONS 10-100 TIMES FASTER THAN VERTICAL*

- *TANGENTIAL FORCES DOMINATE PLATE TECTONICS*

Forces acting on the lithosphere

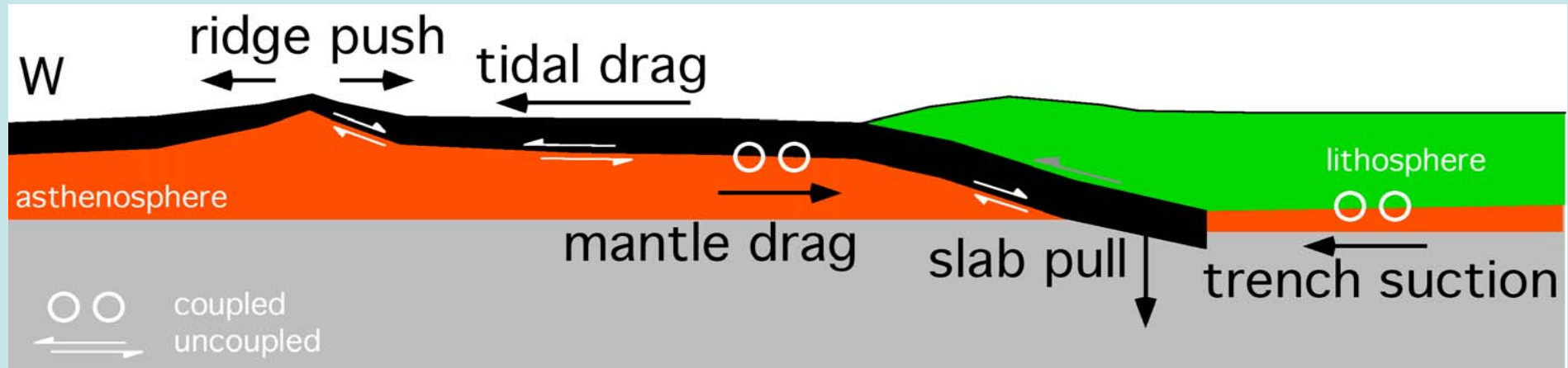
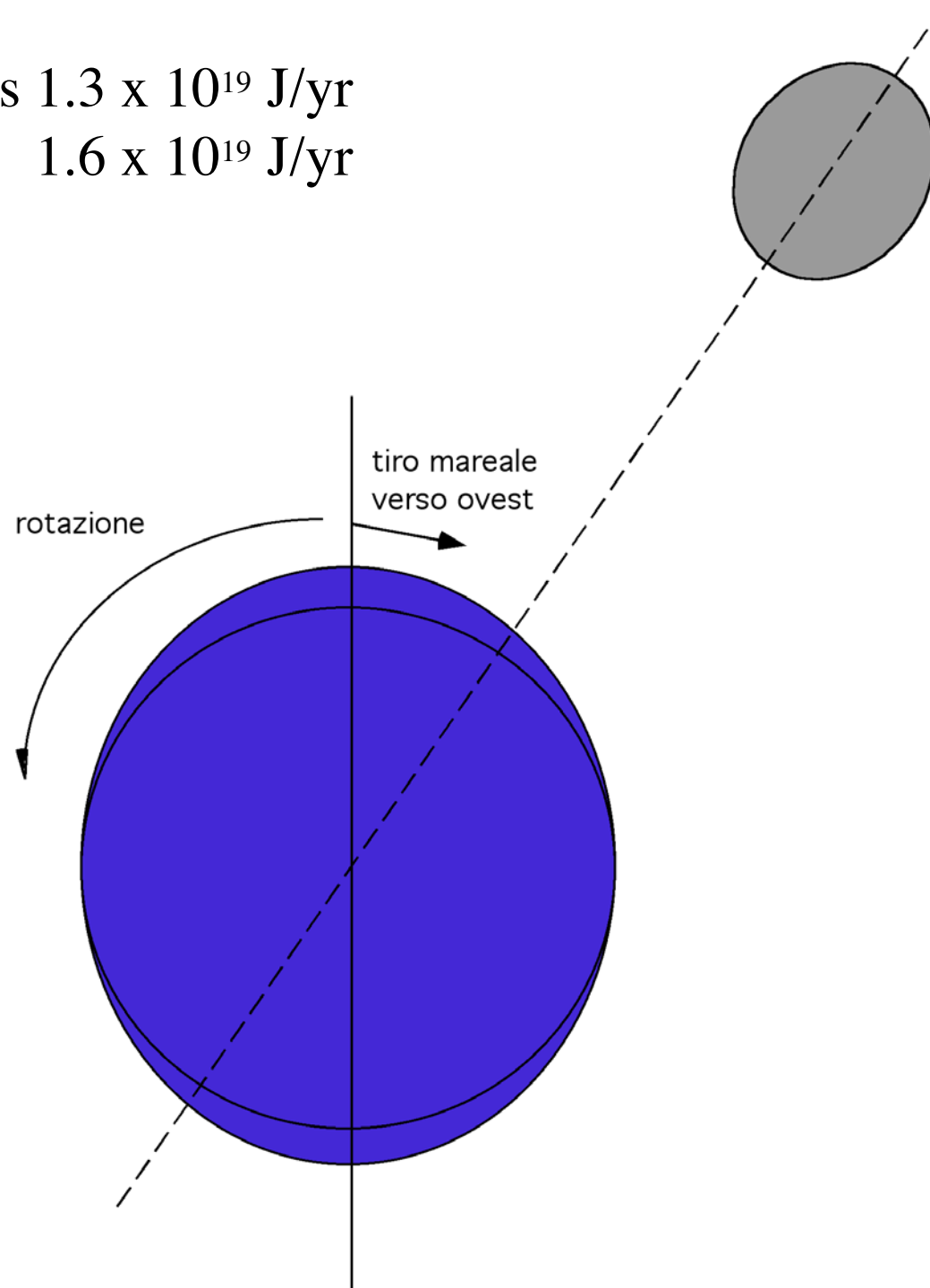
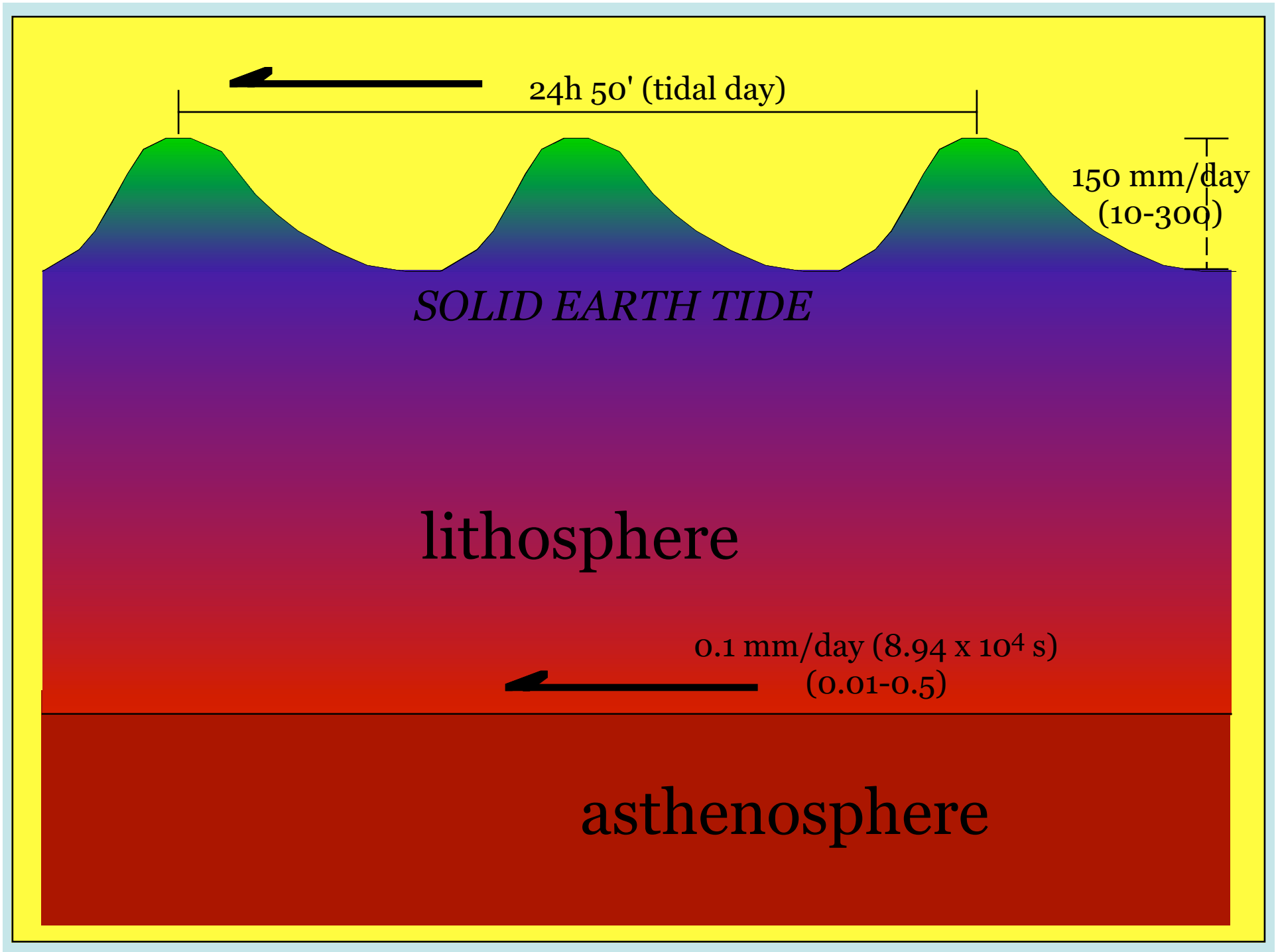
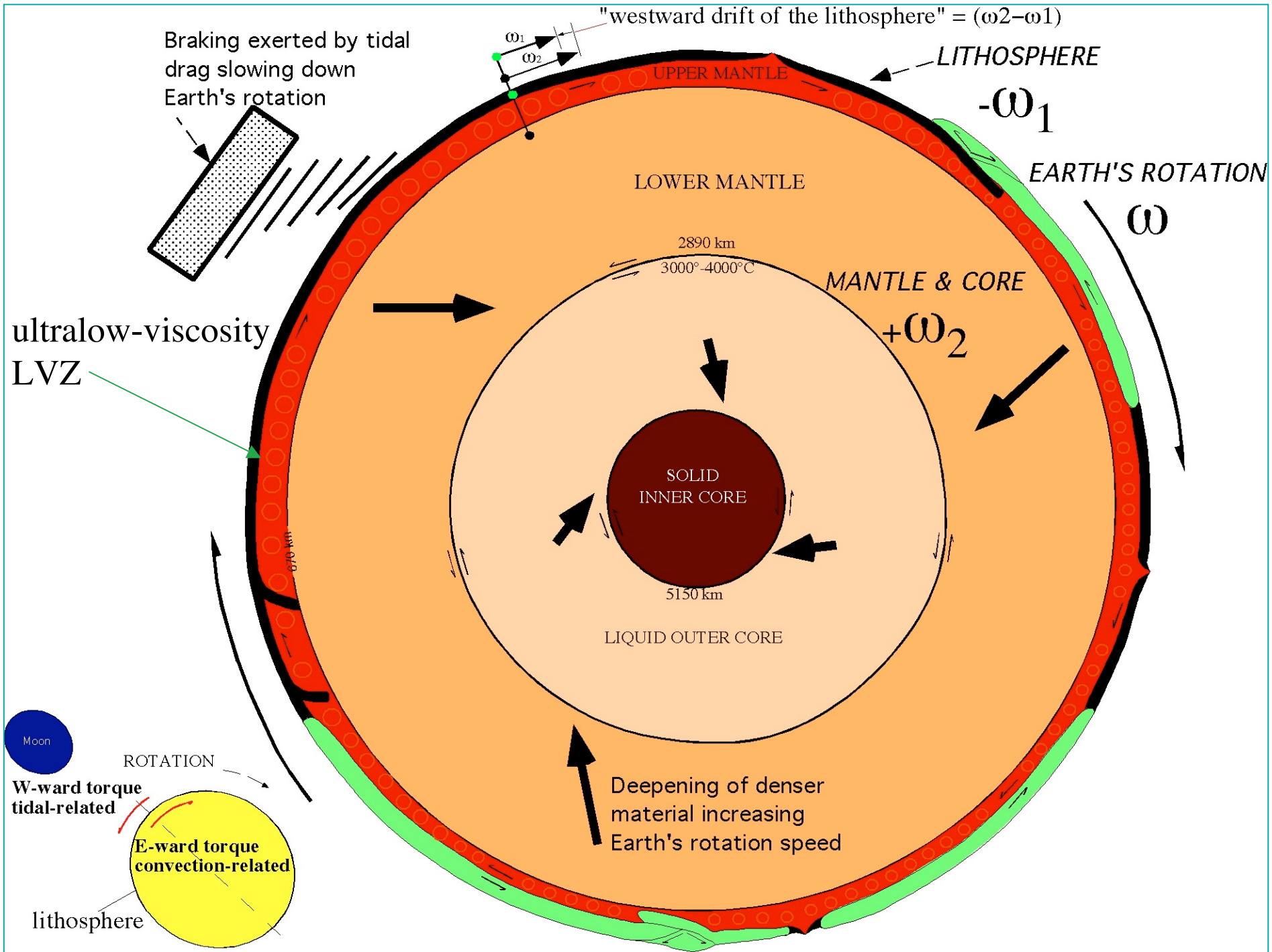


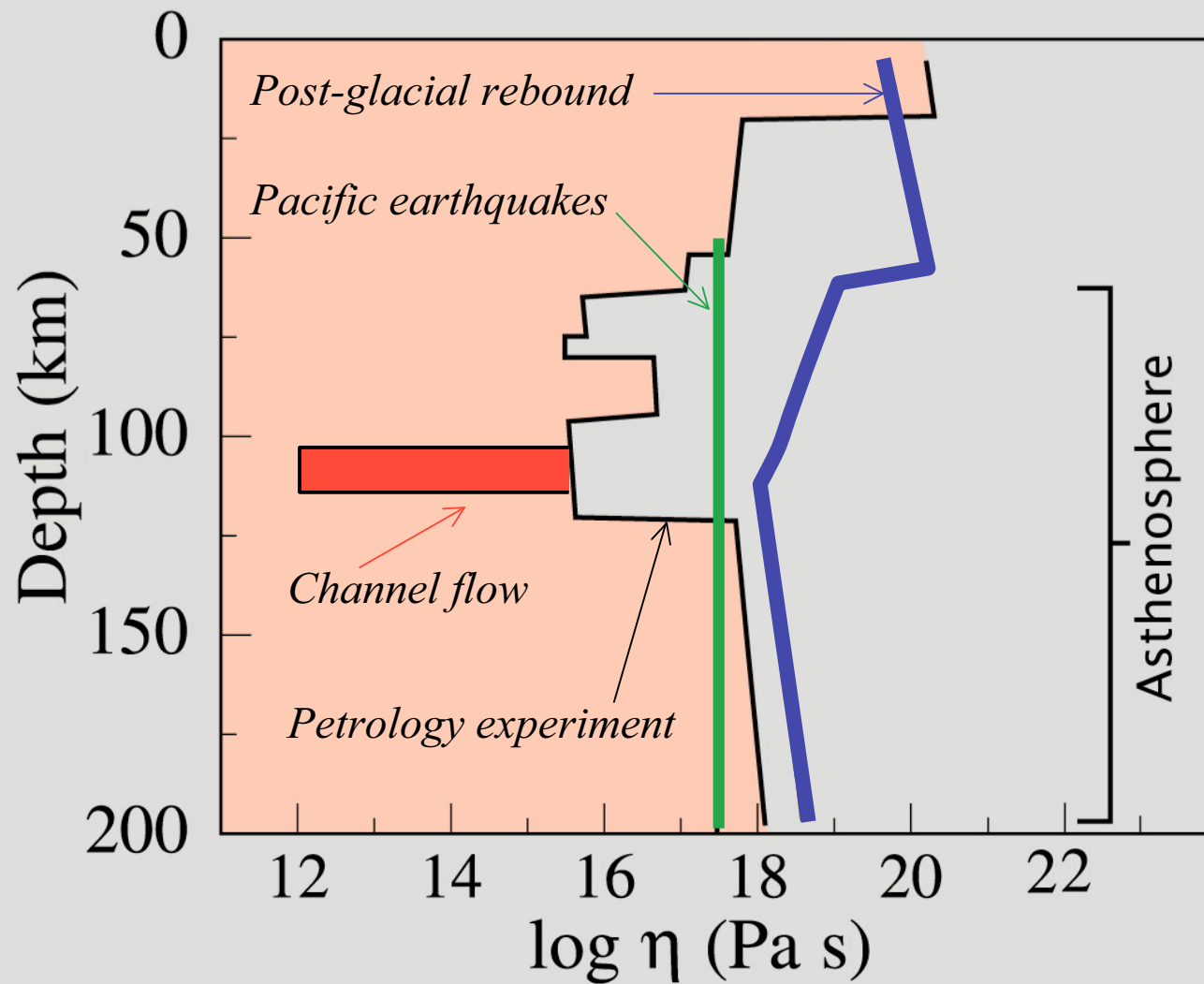
Plate tectonics 1.3×10^{19} J/yr

Tidal friction 1.6×10^{19} J/yr



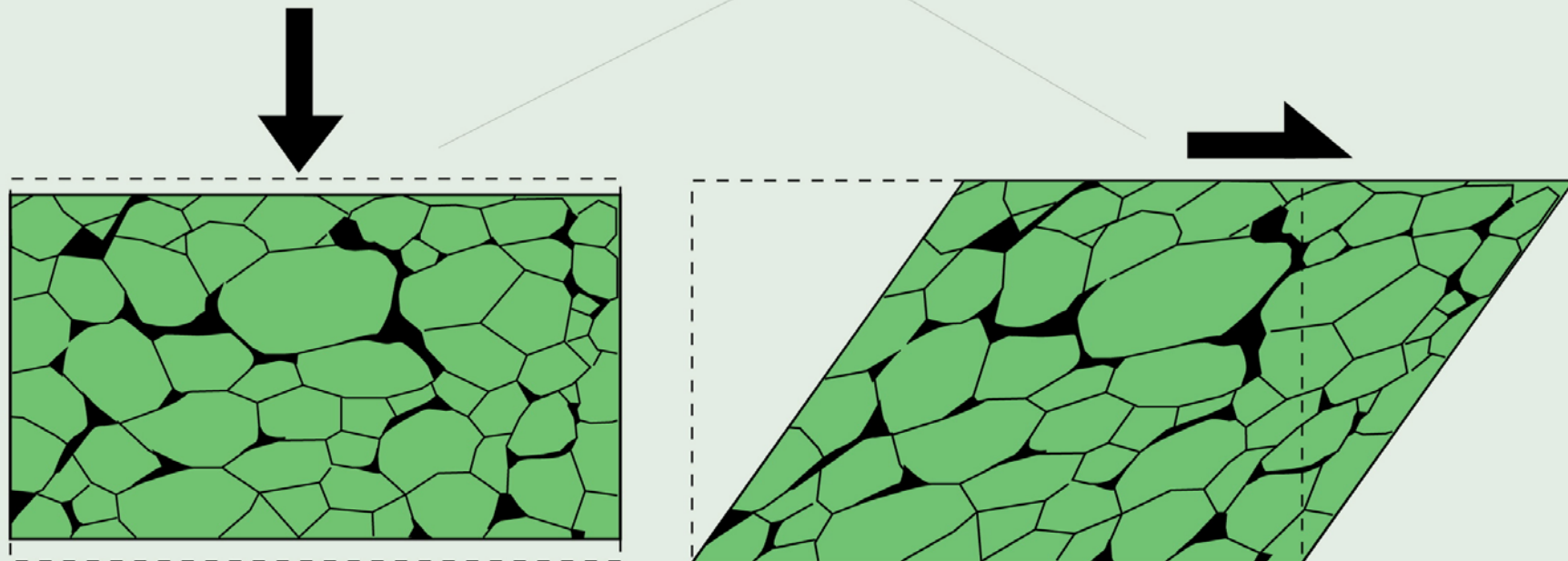
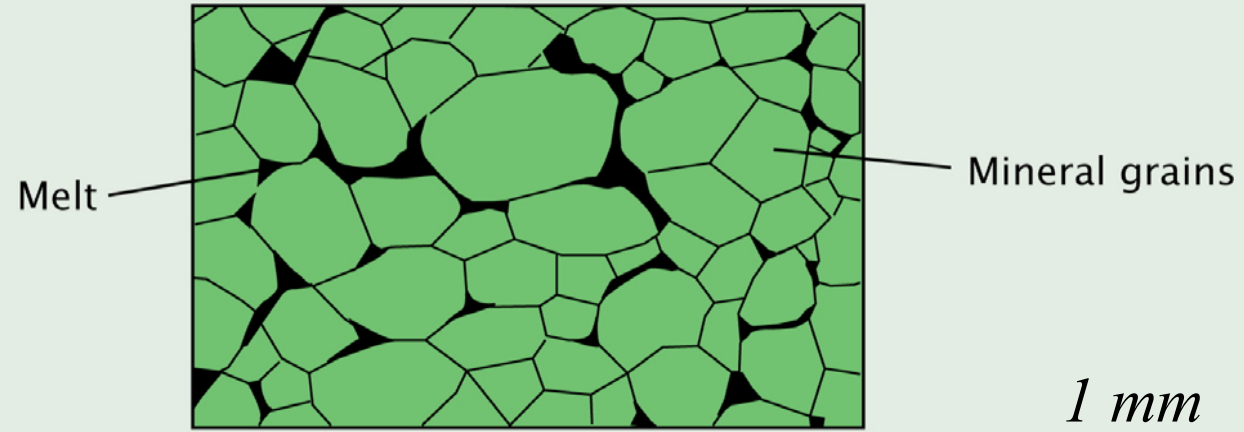




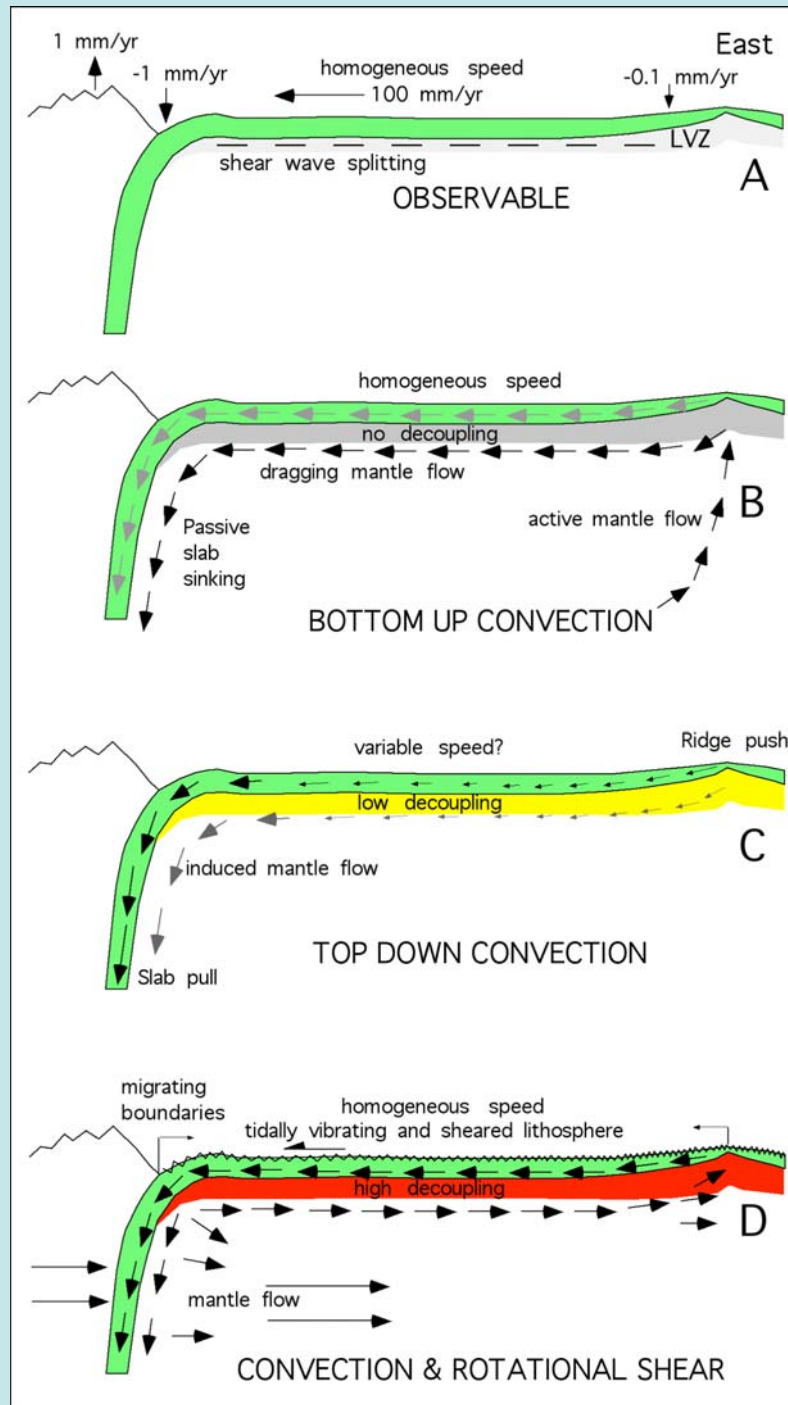


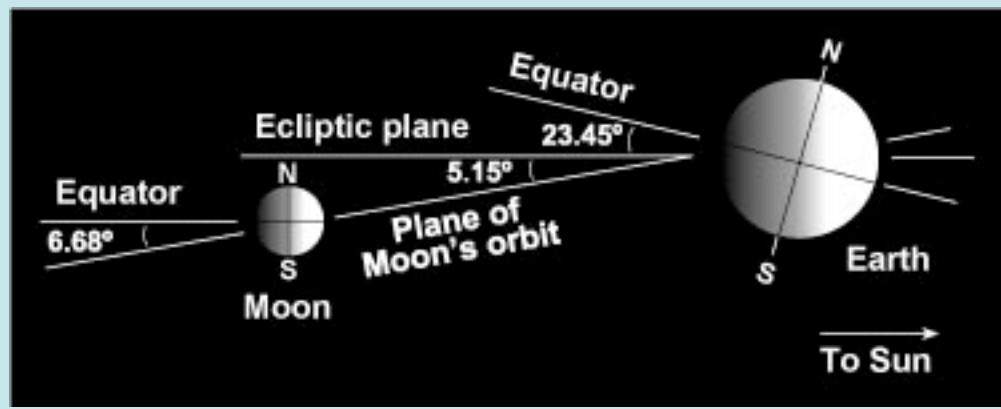
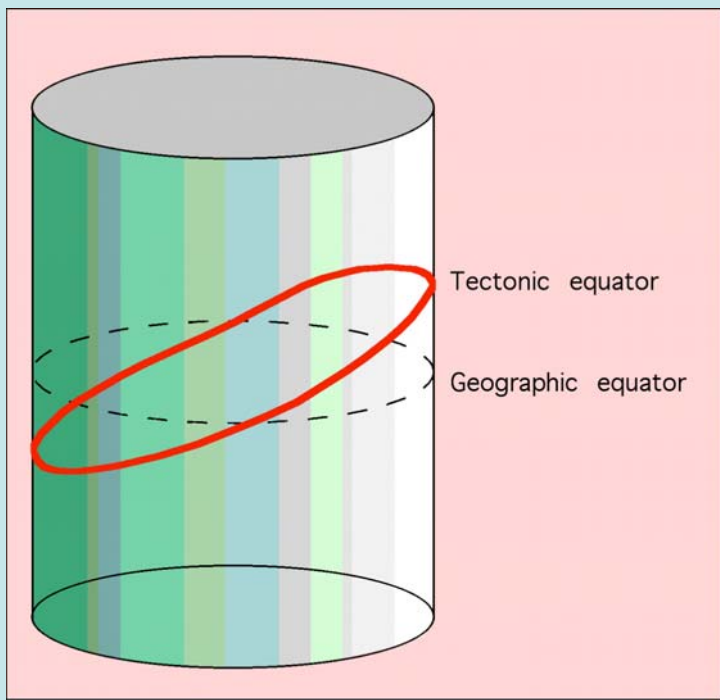
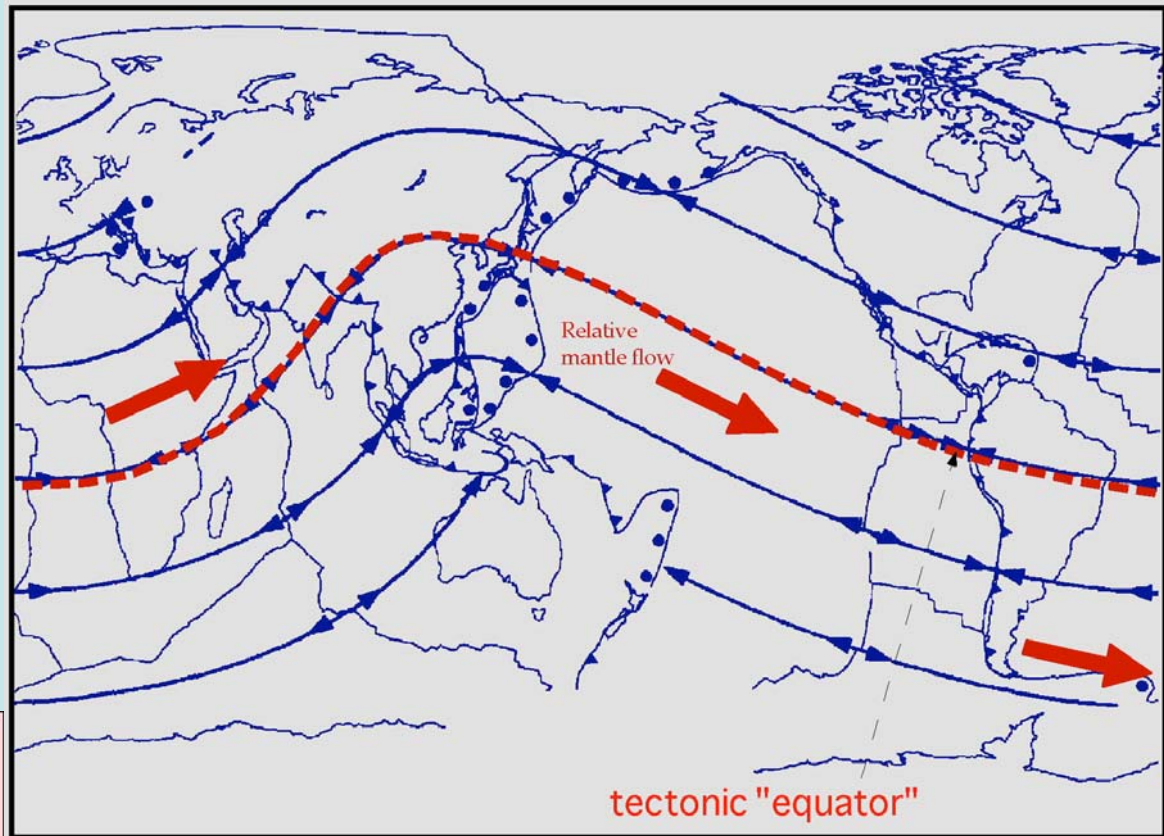
Is the asthenosphere viscosity sufficiently low?

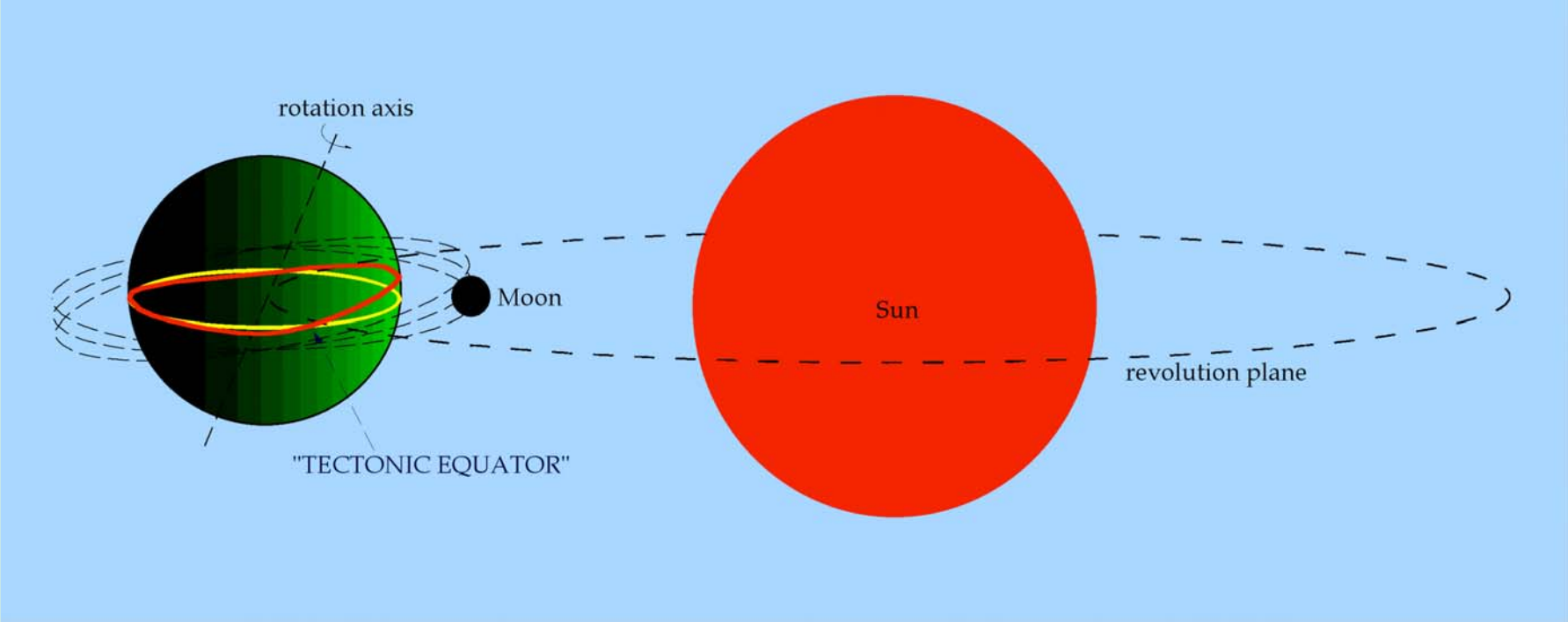
Undeformed partially molten harzburgite



$\uparrow \eta_{\text{vertical}} \gg \eta_{\text{horizontal}}$







EUROPE is a Jupiter's satellite

Rotation period 3d 13h 14.6m

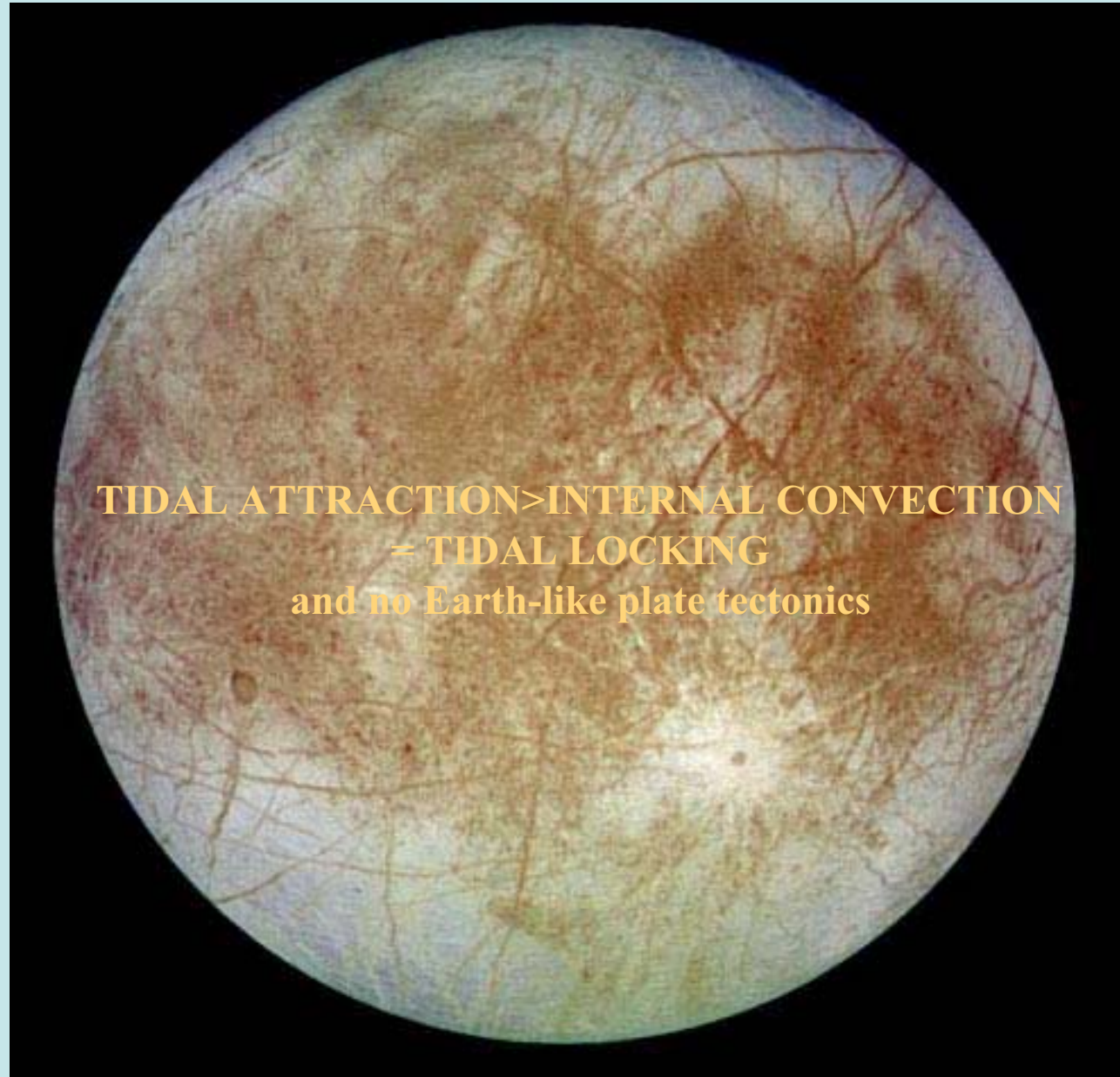
Revolution period 3d 13h 14.6m

(RESONANCE)

diameter 3,138 km

Mean density 3.01 g/cm³

Surface gravity 1.42 m/s²



- 1) The lithosphere moves along a westerly polarized flow
- 2) There is a global tectonic asymmetry
- 3) Plate boundaries are passive features
- 4) Dominant ductile deformation contained in GPS data
- 5) Plate tectonics is tuned by Earth's rotation

Thanks



M. BABINET PREVÉNU PAR SA PORTIÈRE
DE LA VISITE DE LA COMÈTE

Lithograph. Honoré Daumier, French, 1808-1879

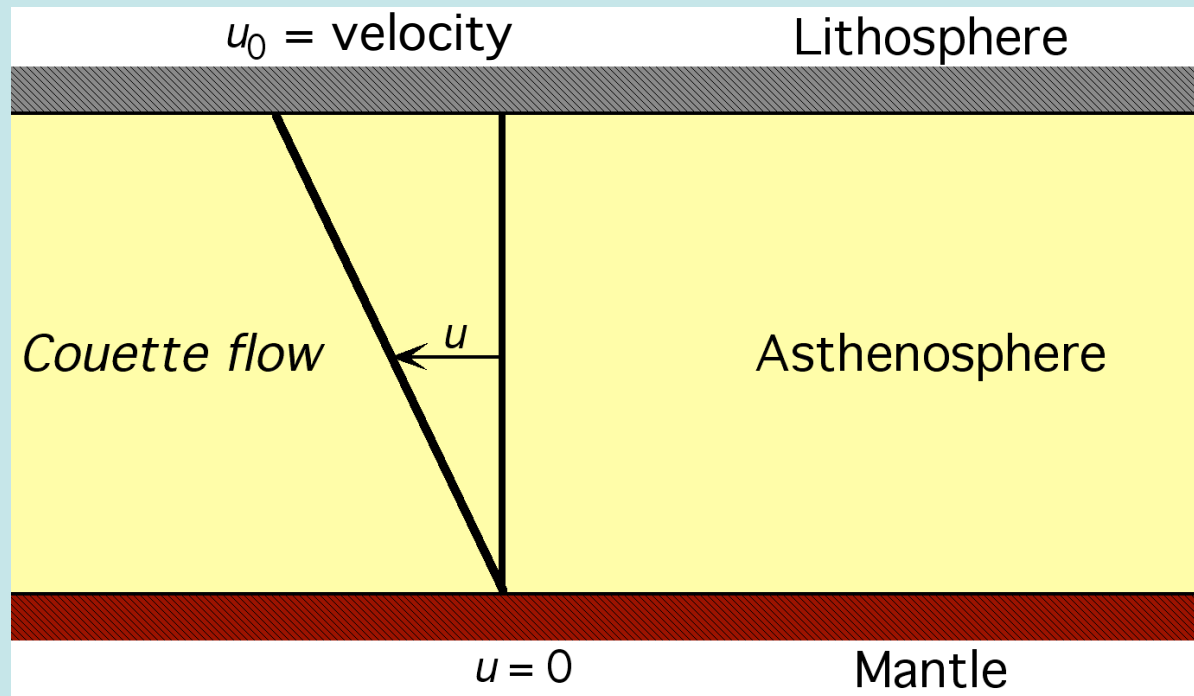
MUSEUM OF FINE ARTS, BOSTON

The maximum excess temperature due to frictional heating with respect to the temperature linear profile in a Couette flow is:

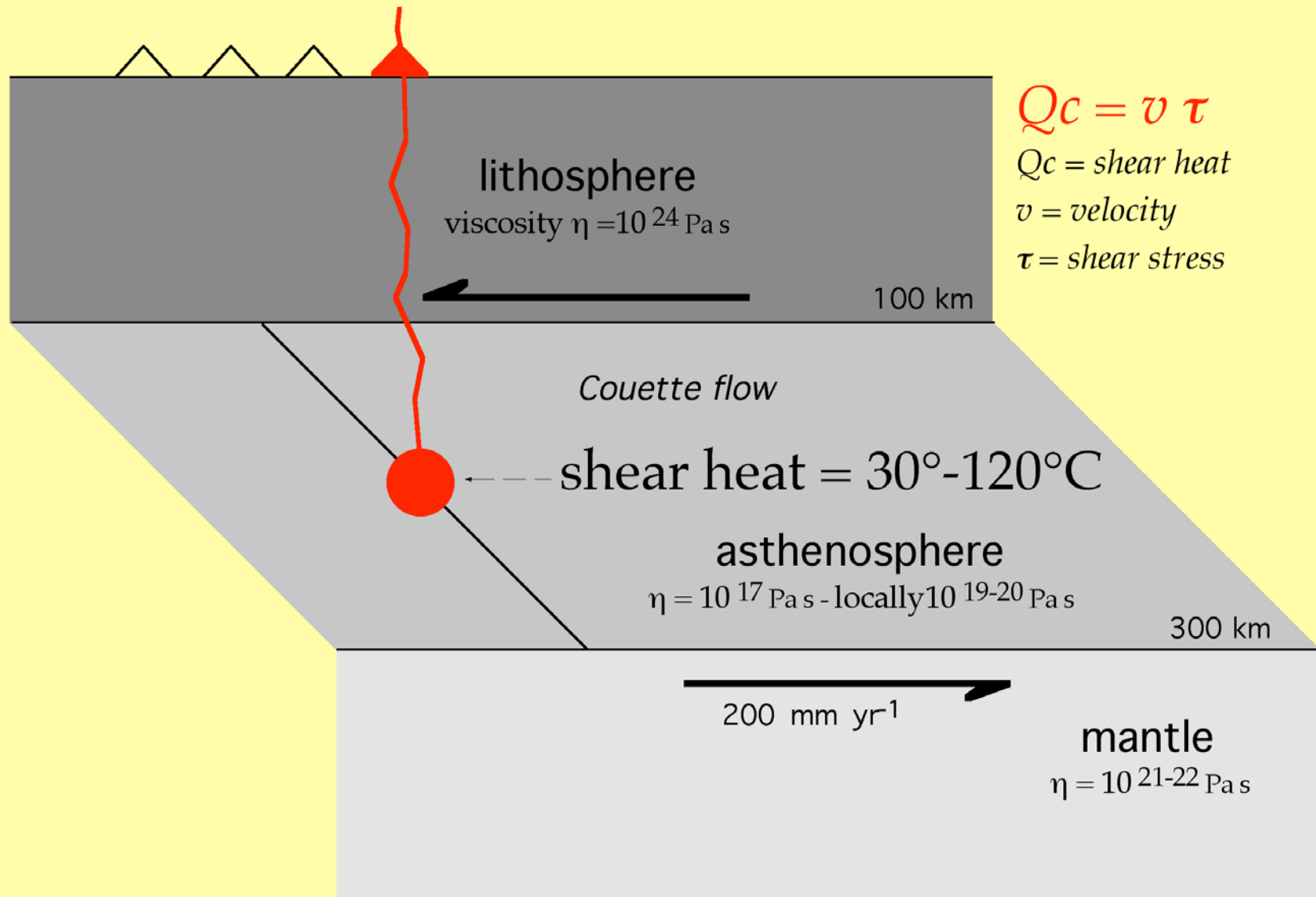
$$\theta_e^{max} = \frac{Pr E}{8}$$

and occurs in the middle of the flow

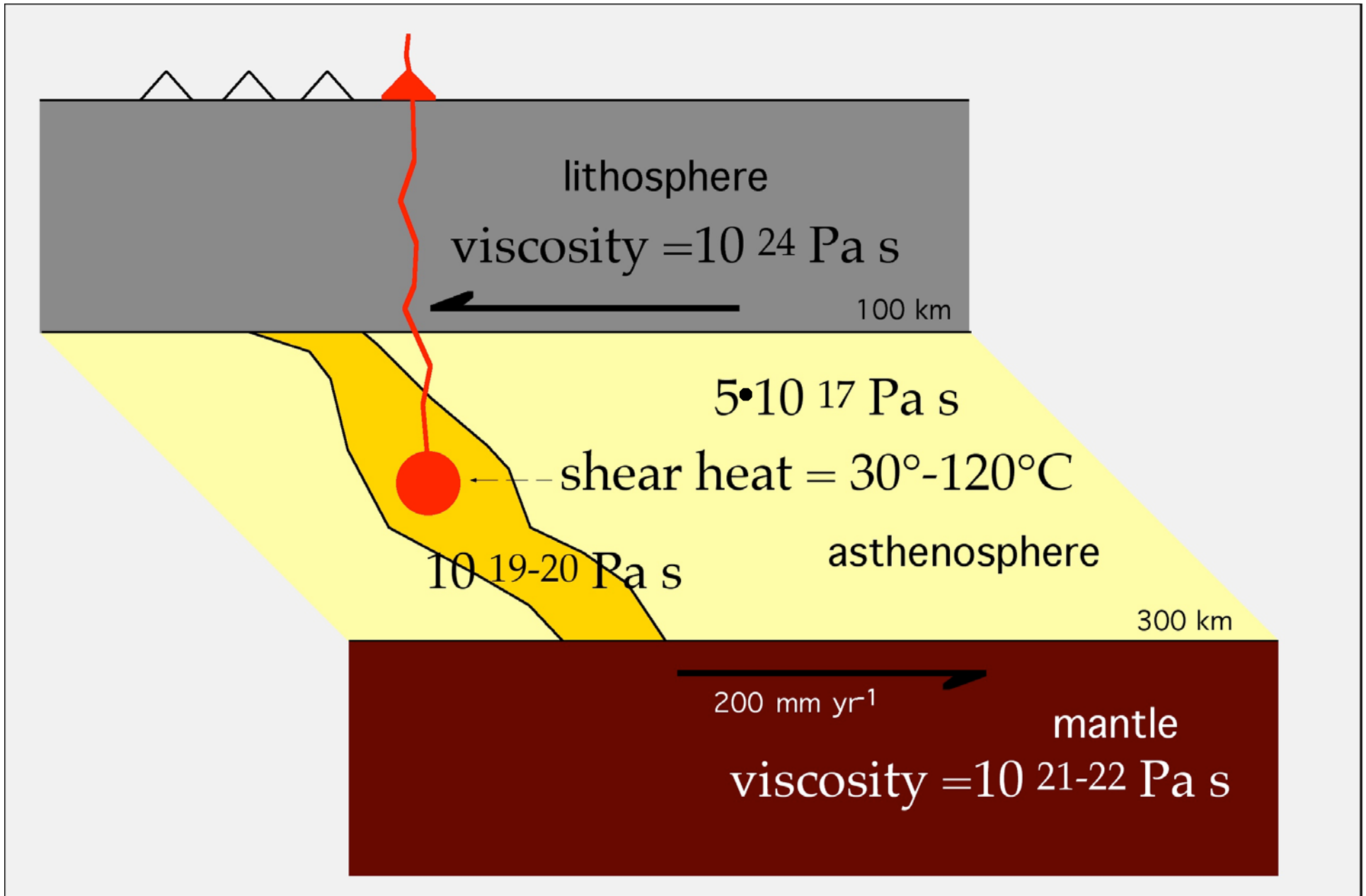
where E is the Eckert number and Pr is the Prandtl number (Turcotte and Schubert, 2002)



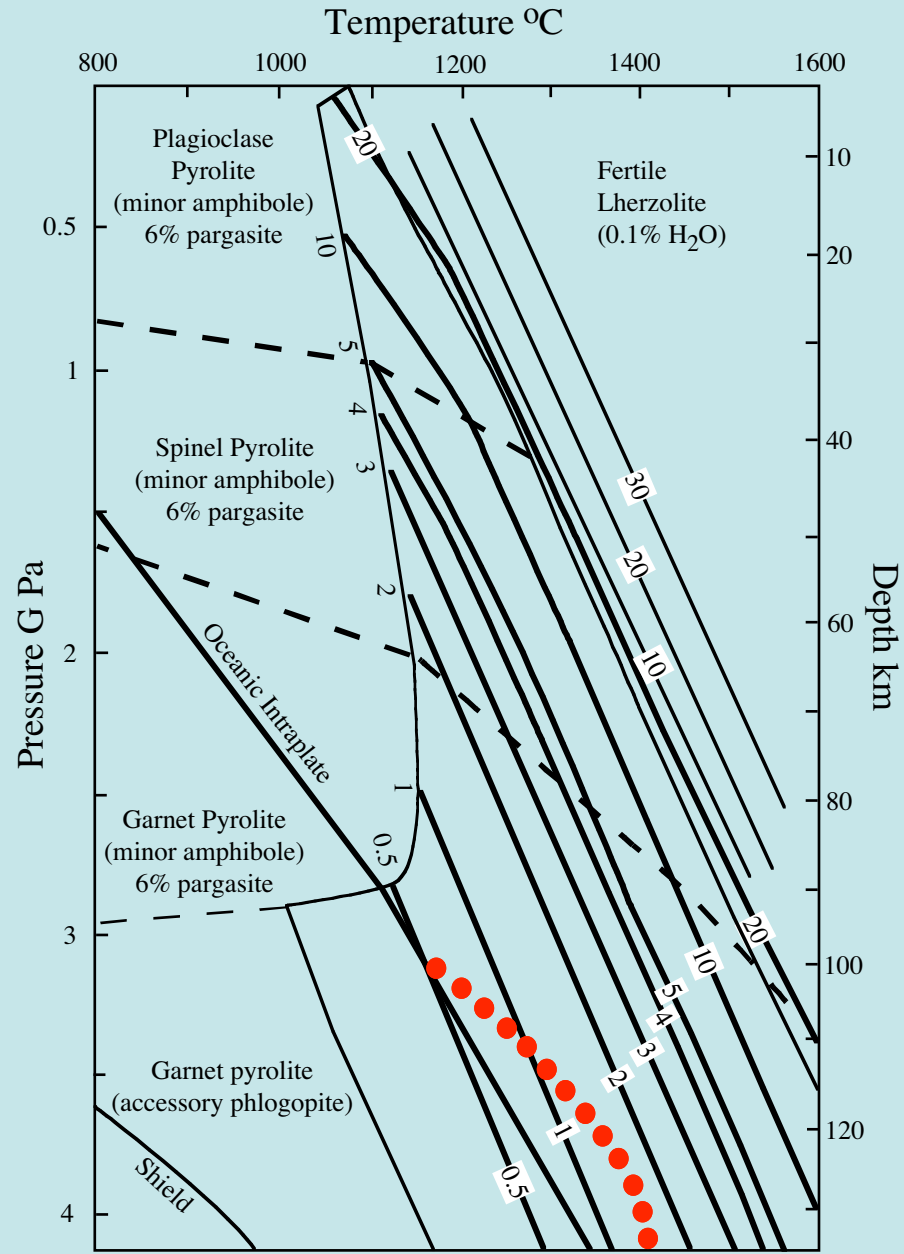
Dogliani, Green and Mongelli, 2005

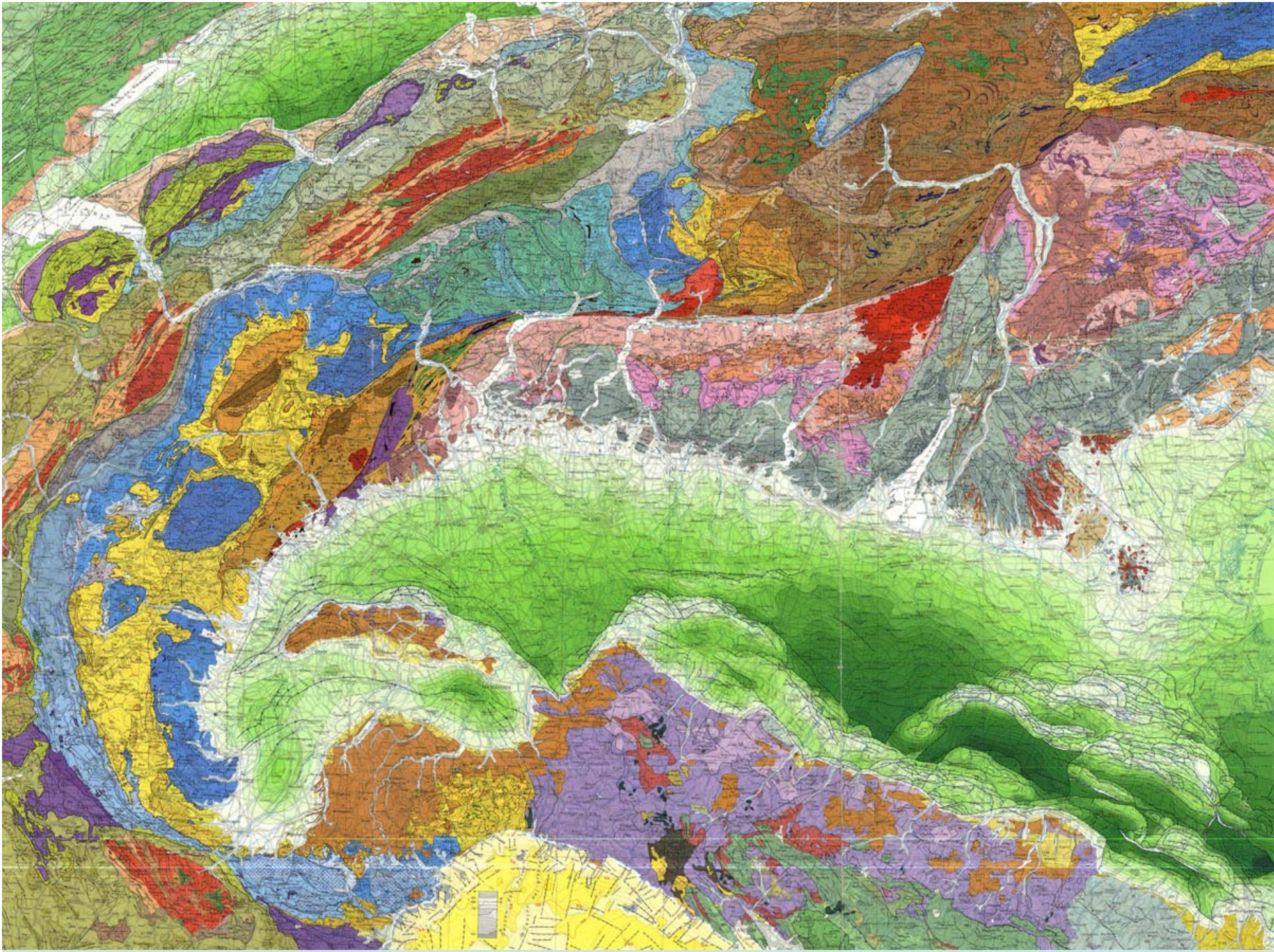


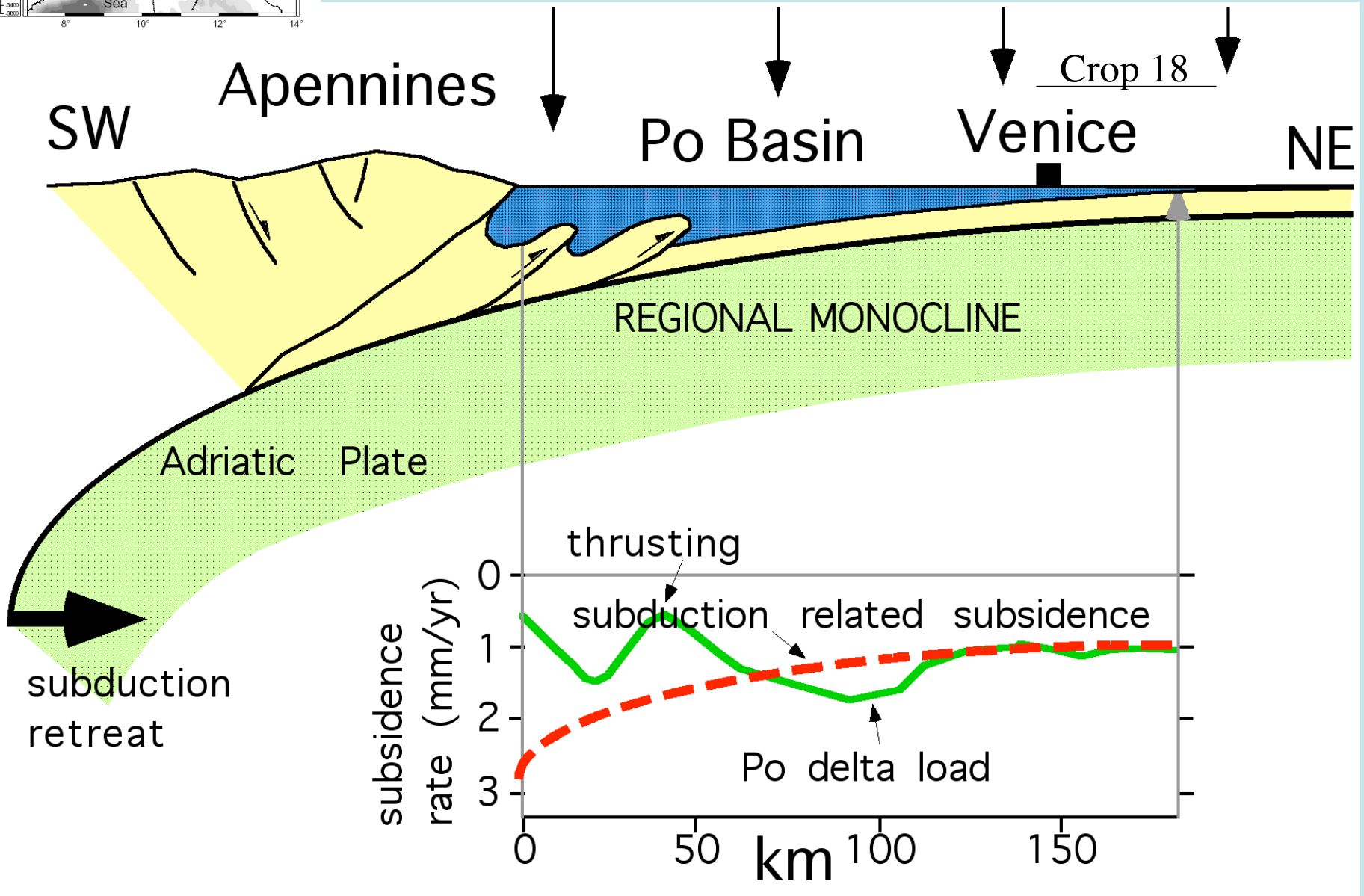
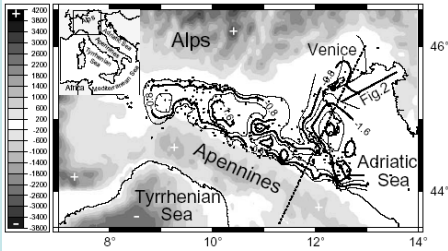
- Local increase of the asthenospheric viscosity can generate extra T (i.e. higher viscosity = larger amount of melting?)



- Local increase of the asthenospheric viscosity can generate extra T (i.e. higher viscosity = larger amount of melting?)



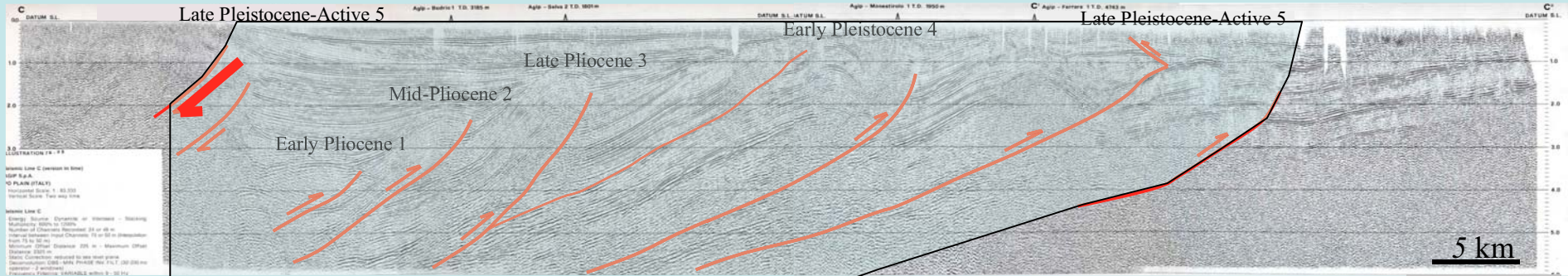




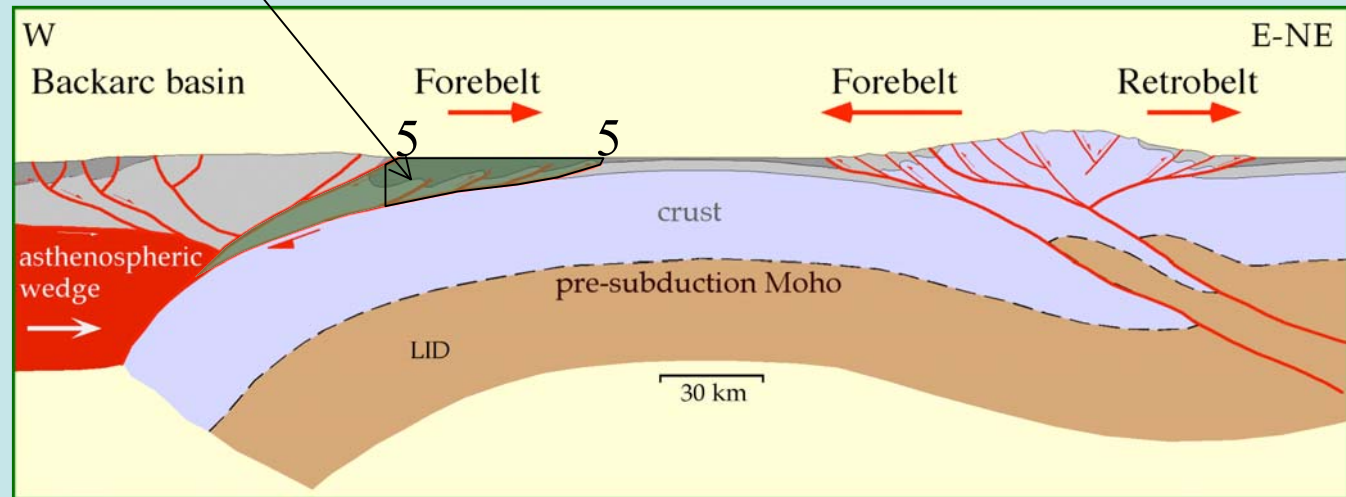
Out-of-sequence thrust
5

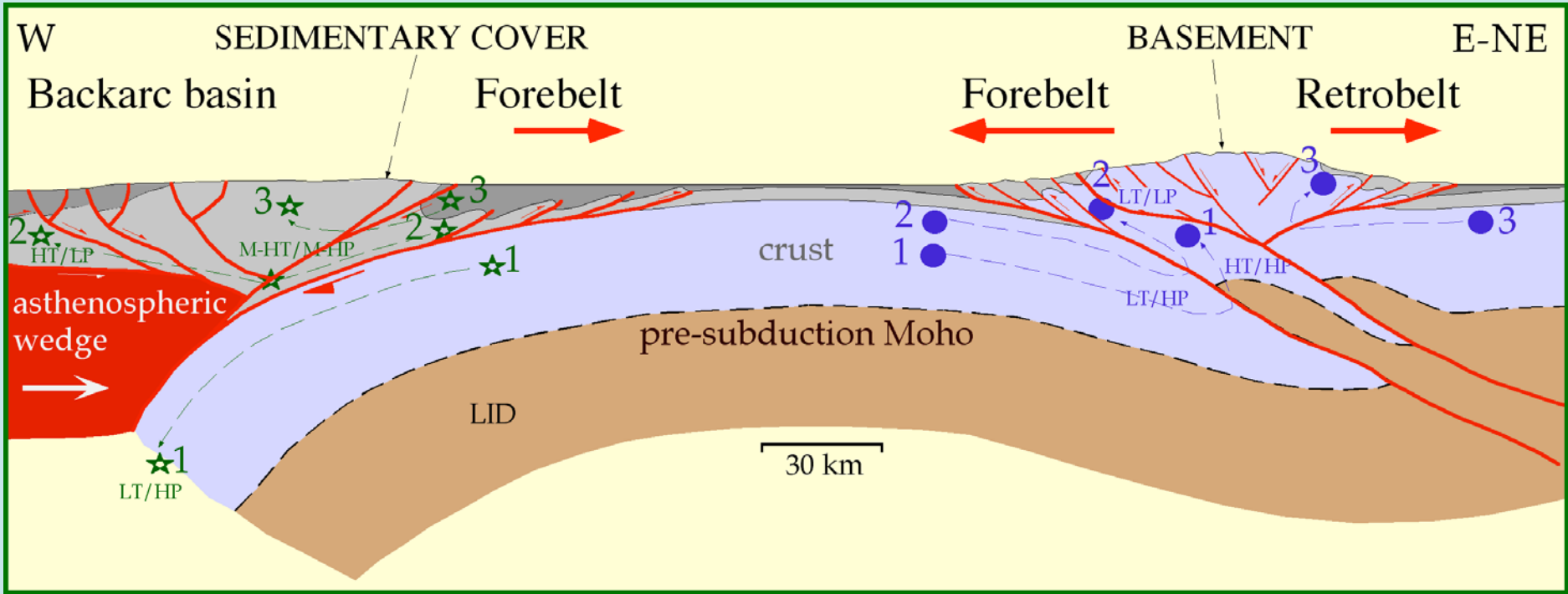
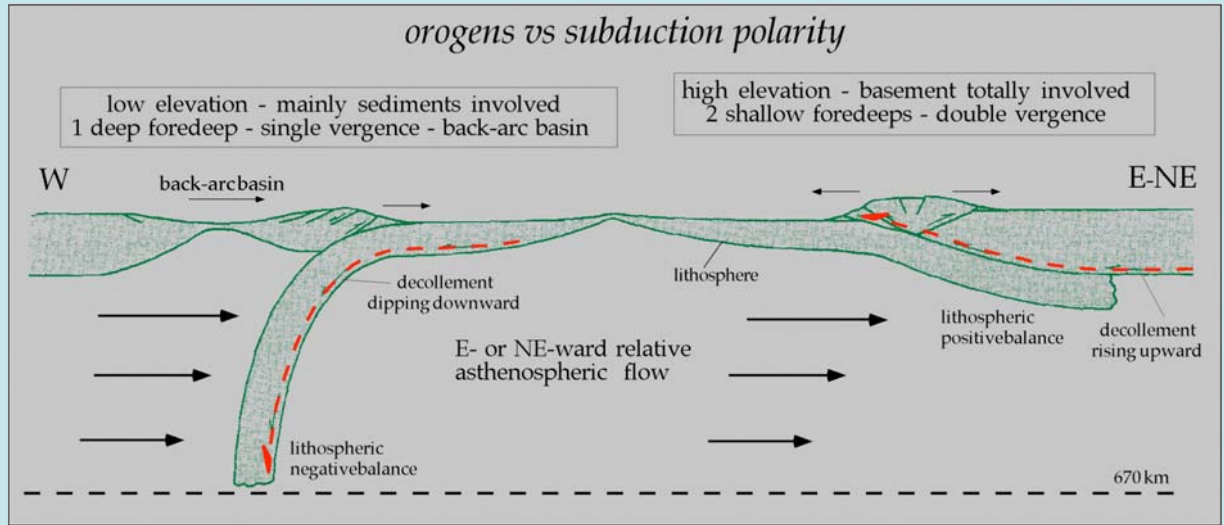
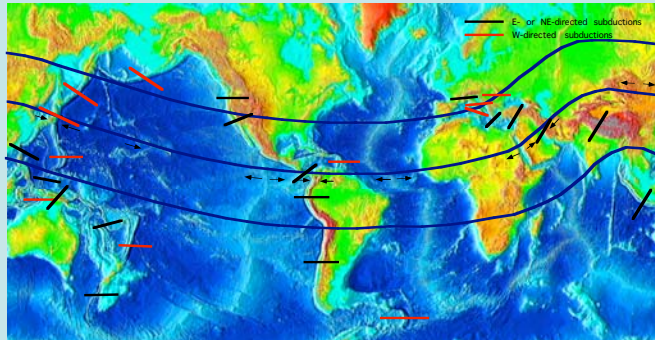
Present *N-Apennines accretionary prism*

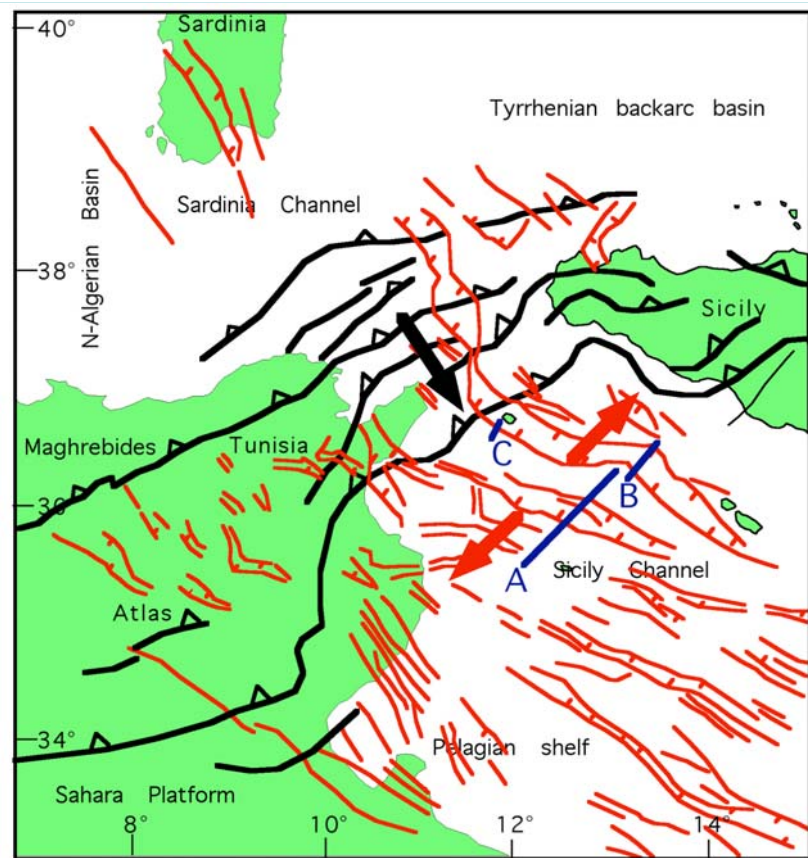
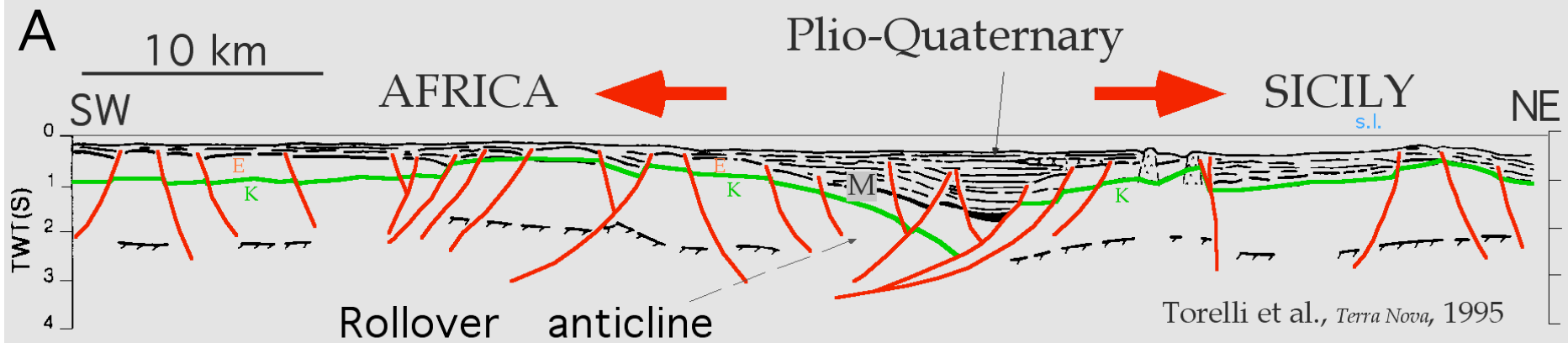
Frontal thrust
5



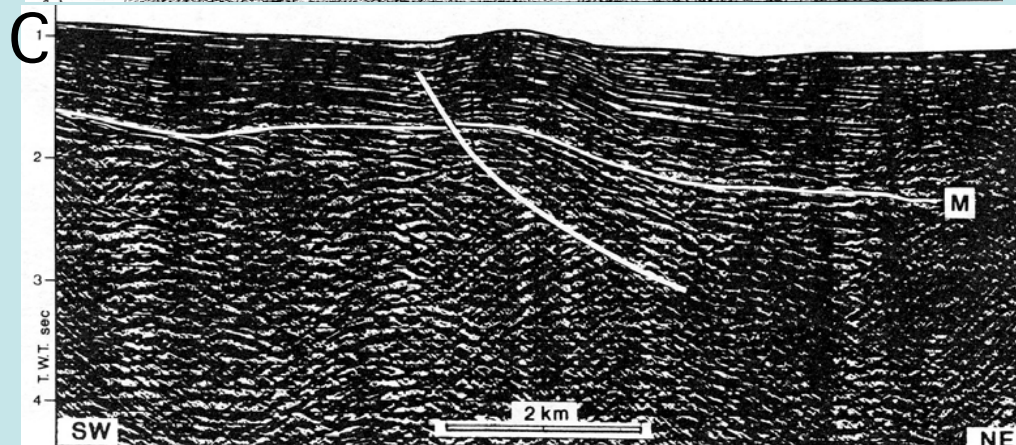
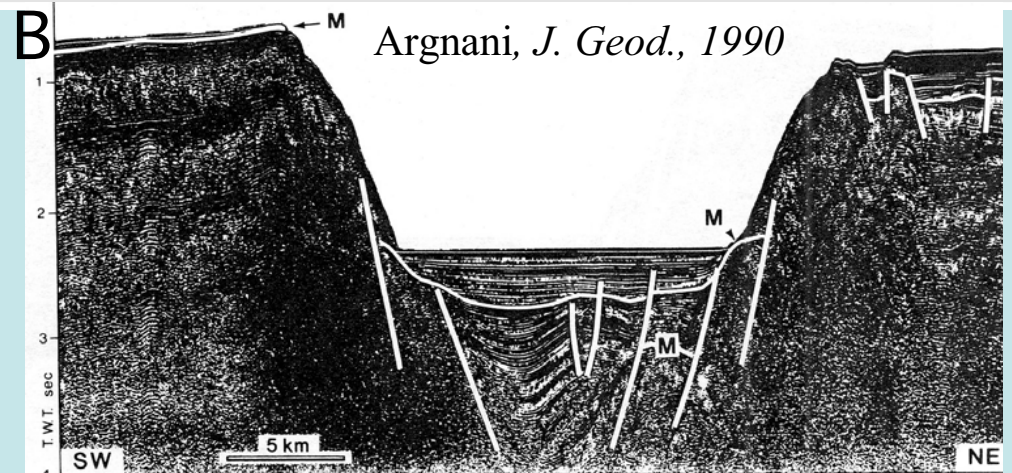
Po Basin - seismics after Pieri, 1983

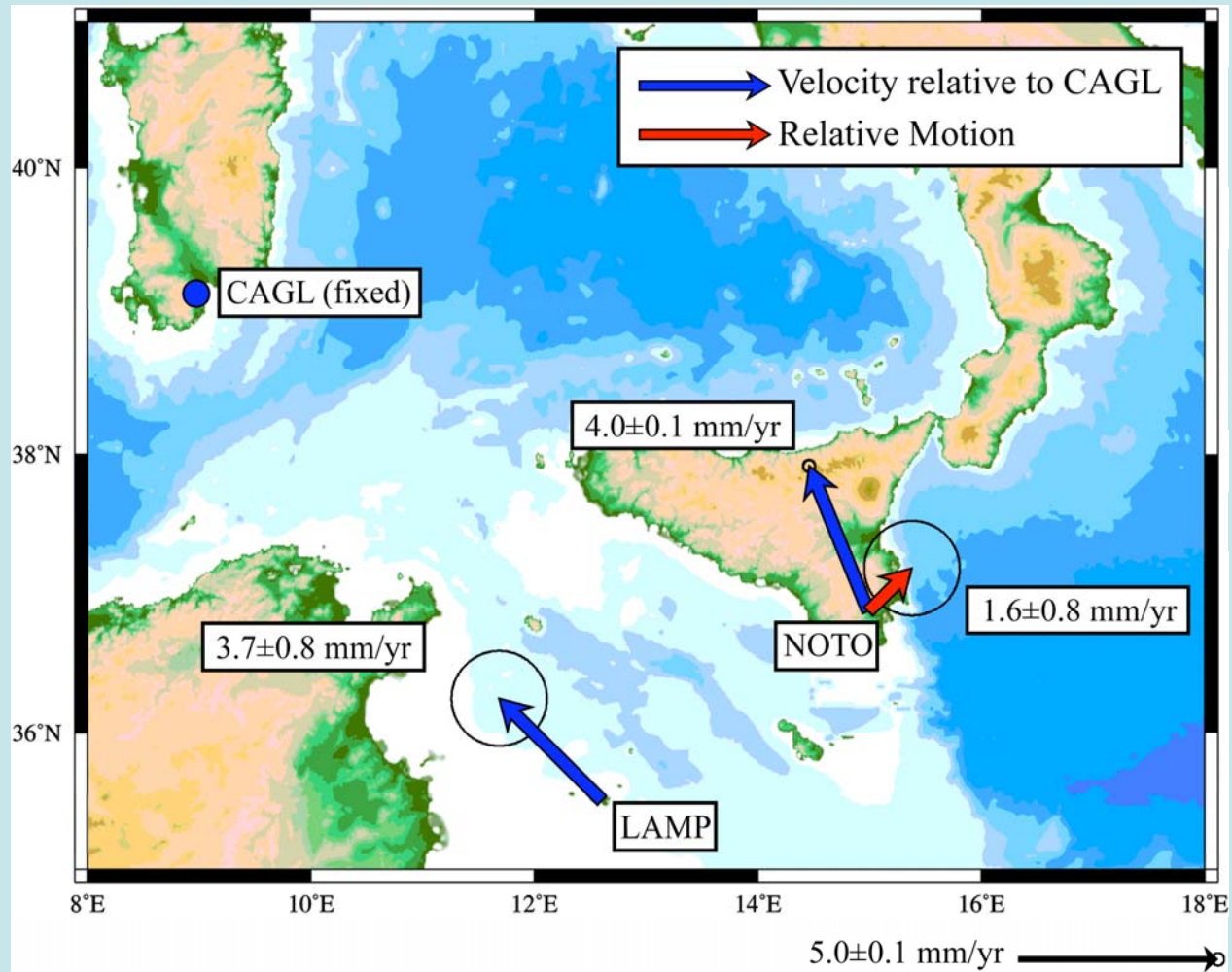
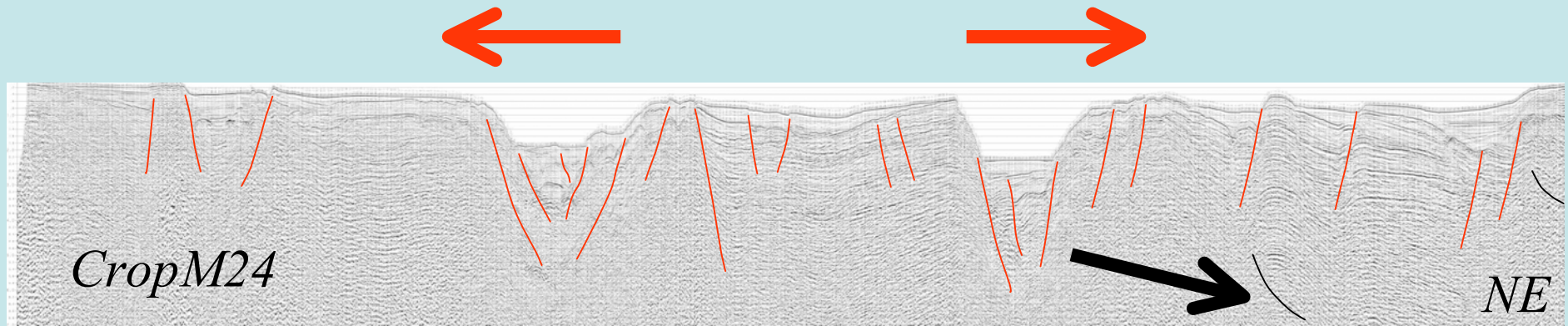






Tricart et al., *Tectonophysics*, 1994

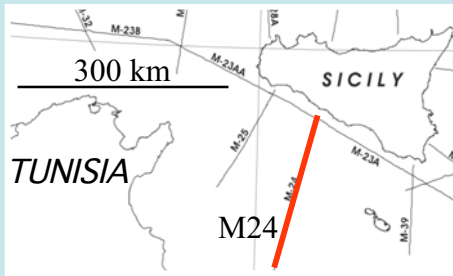
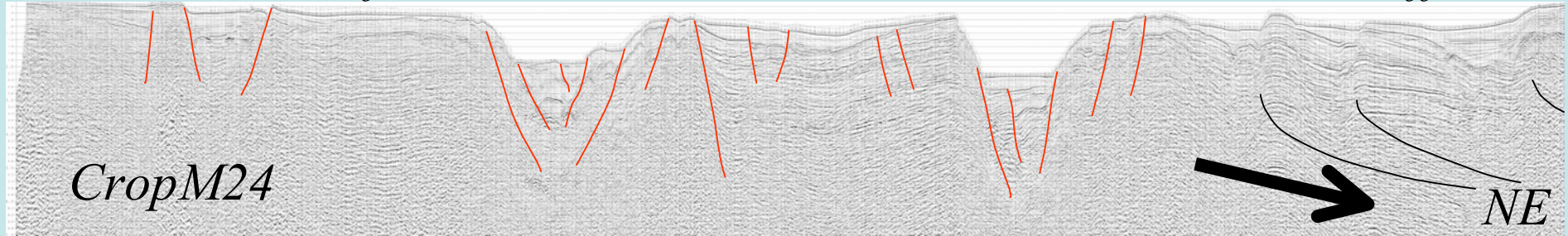




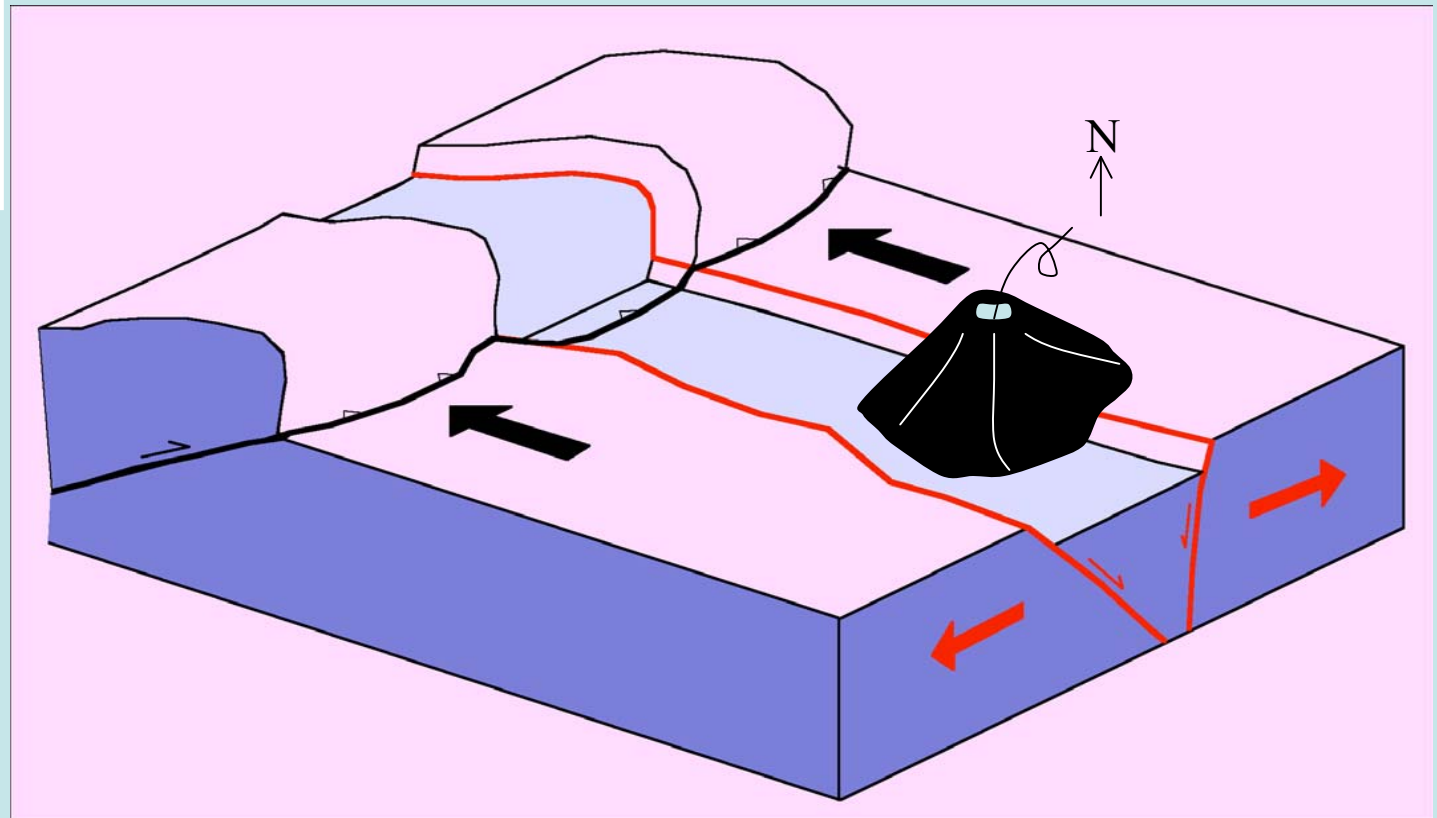
PELAGIAN shelf

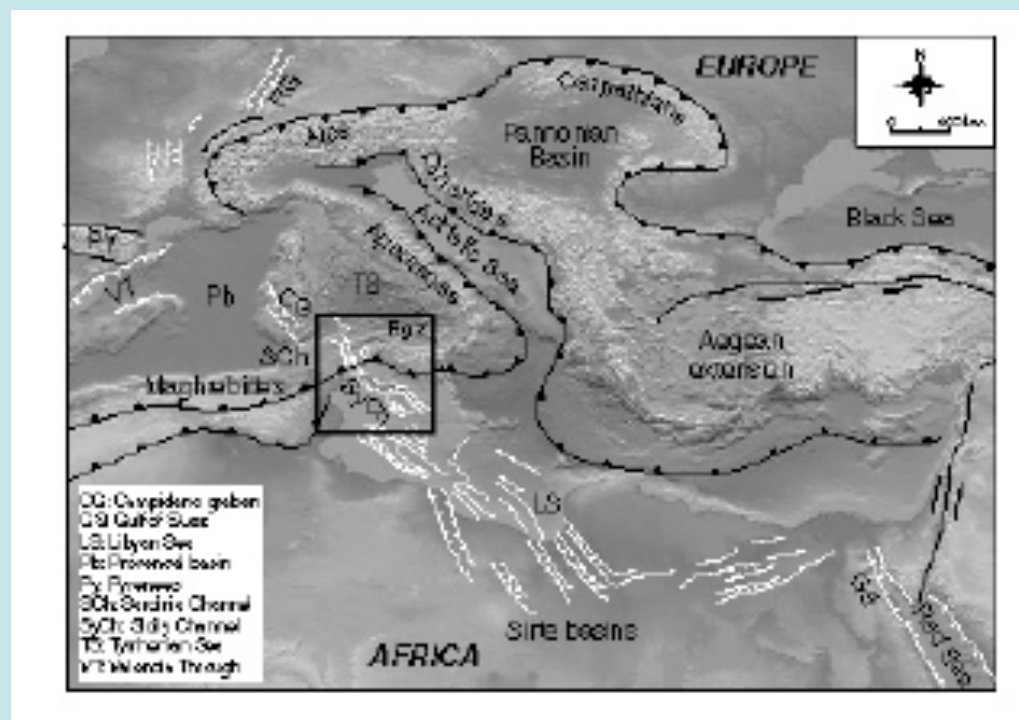
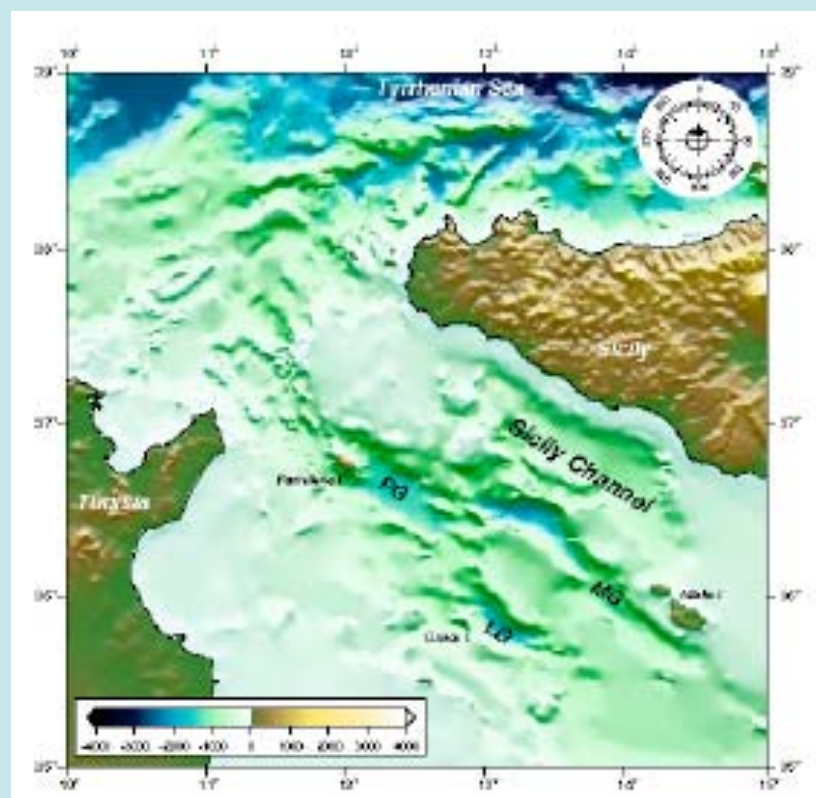


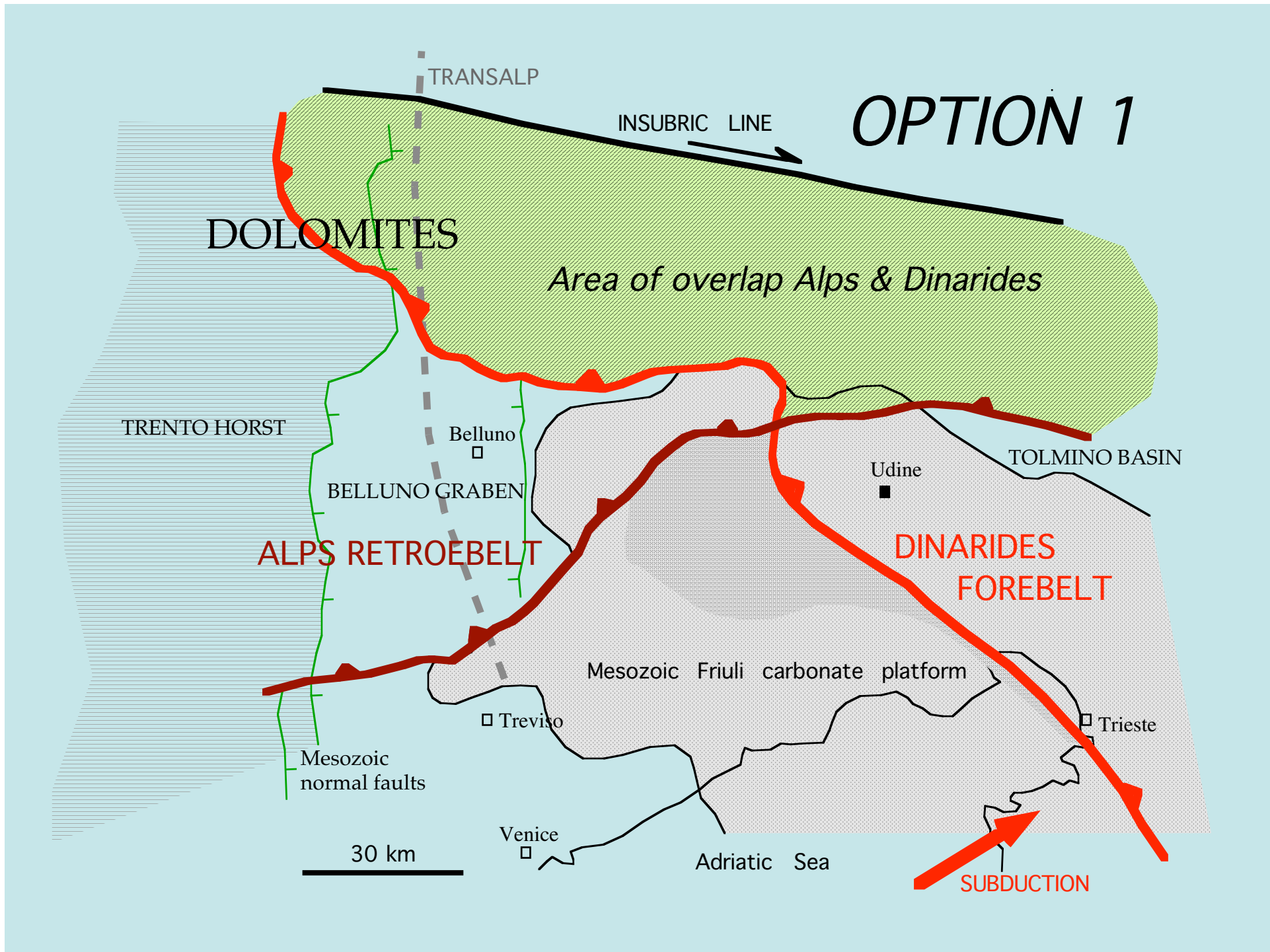
SICILY offshore

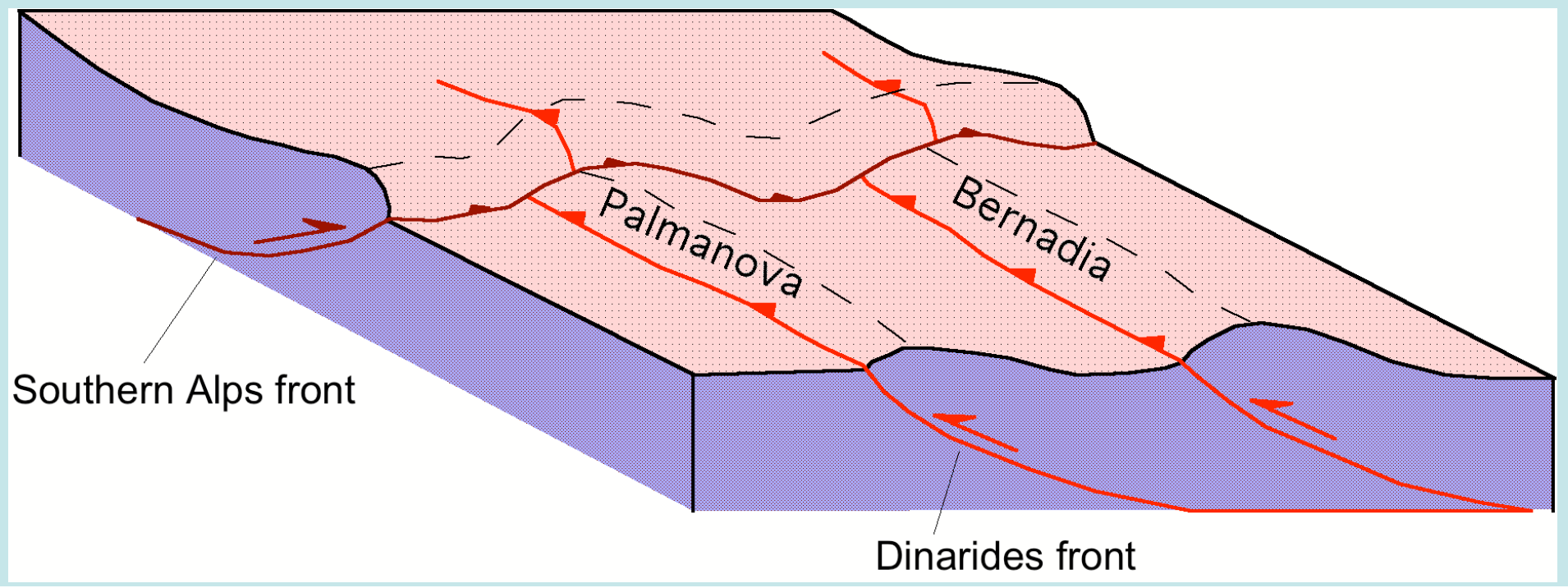
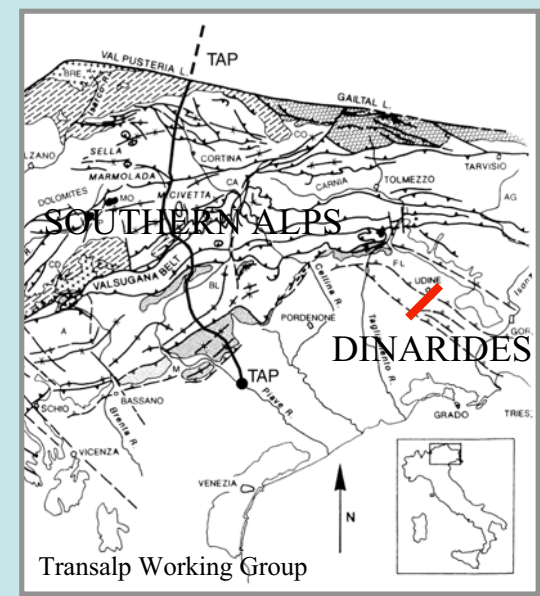
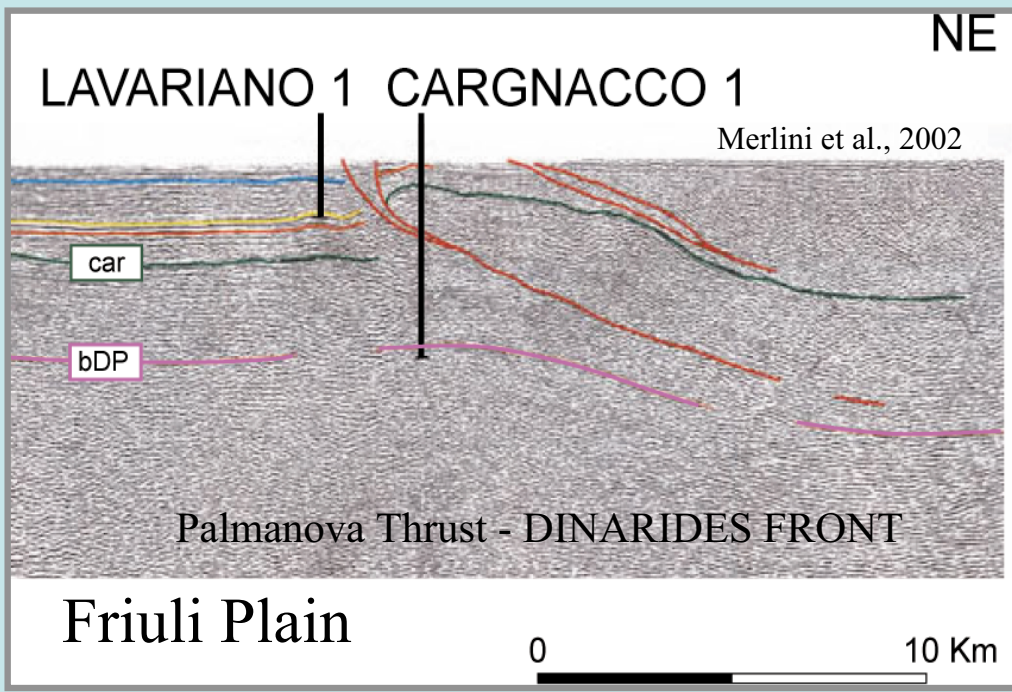


Coexisting
tectonic fields

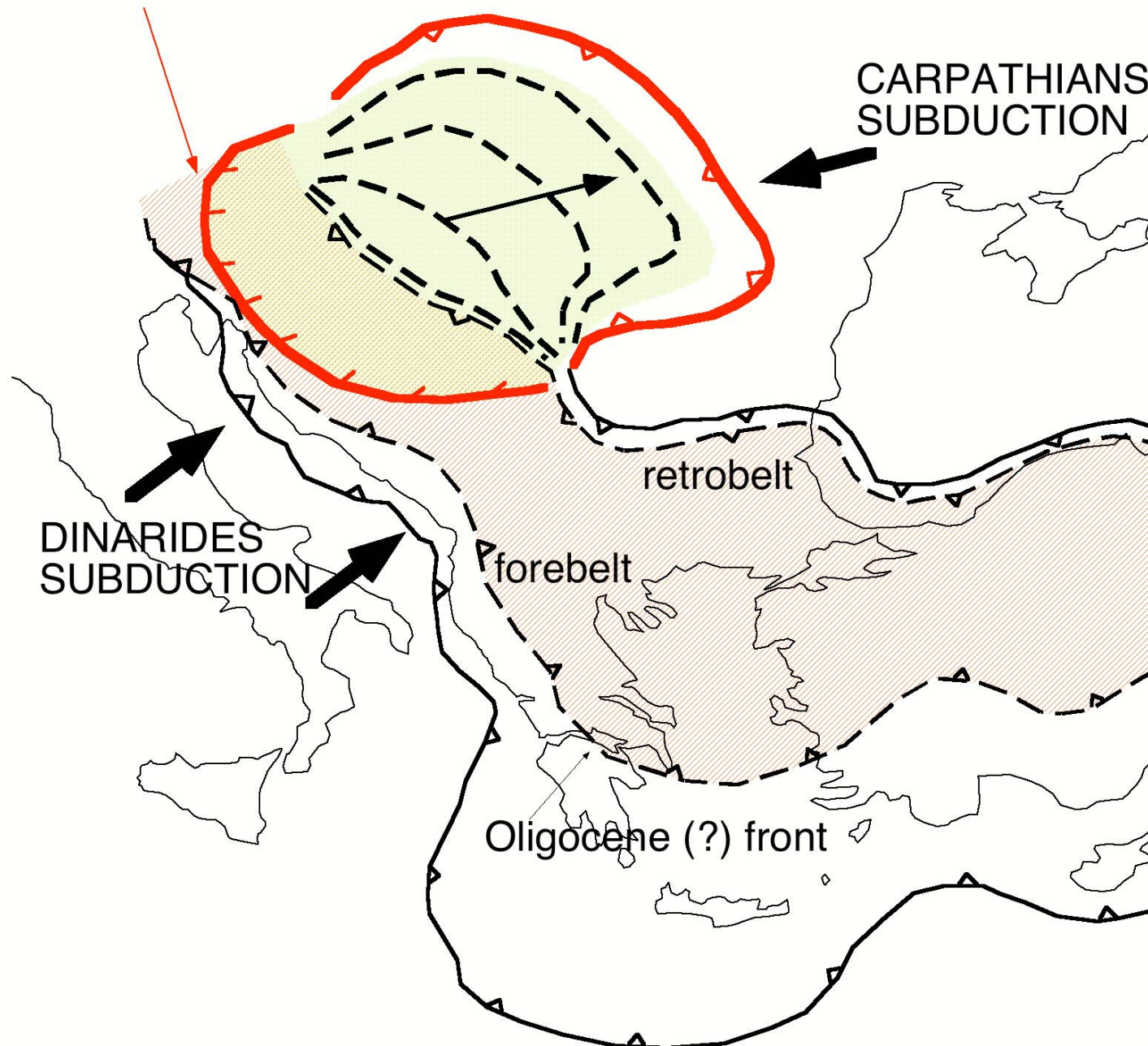


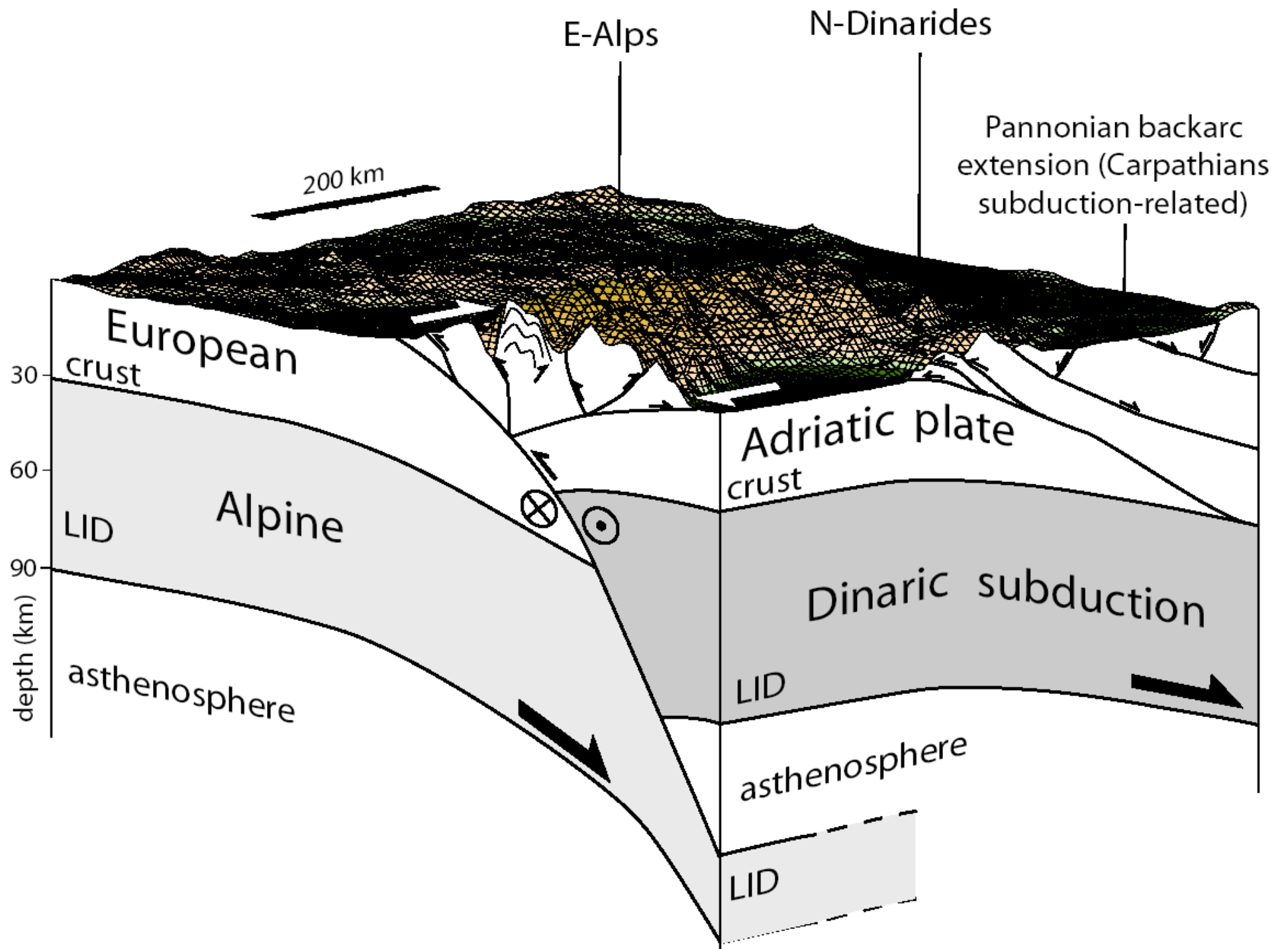






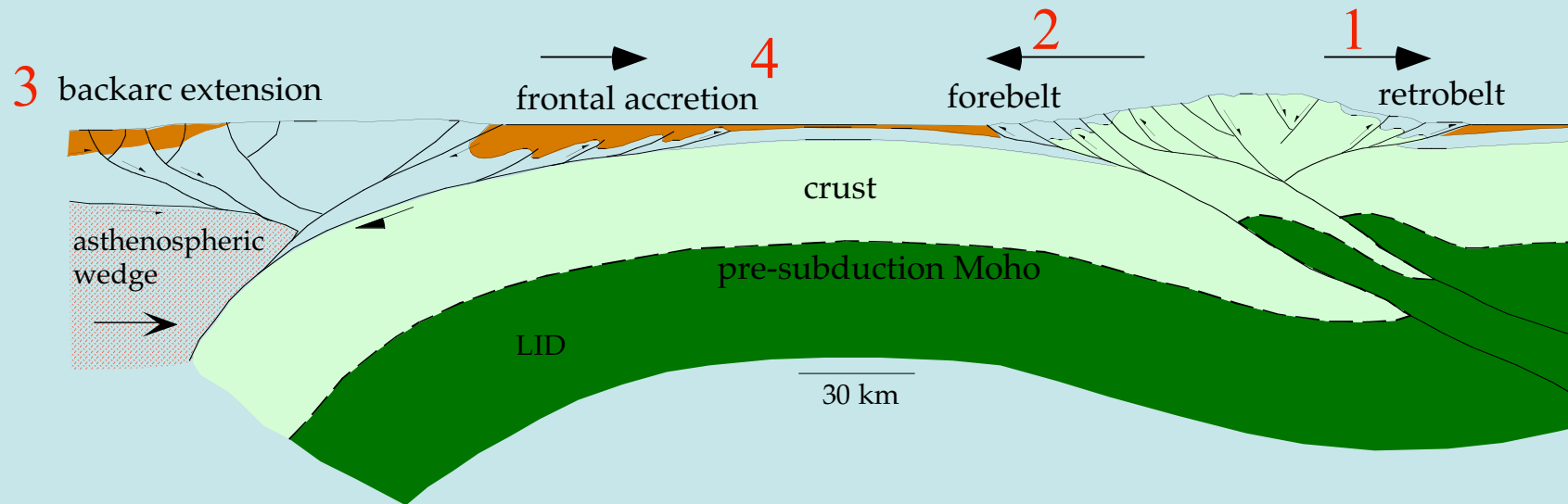
Pannonian extension western margin



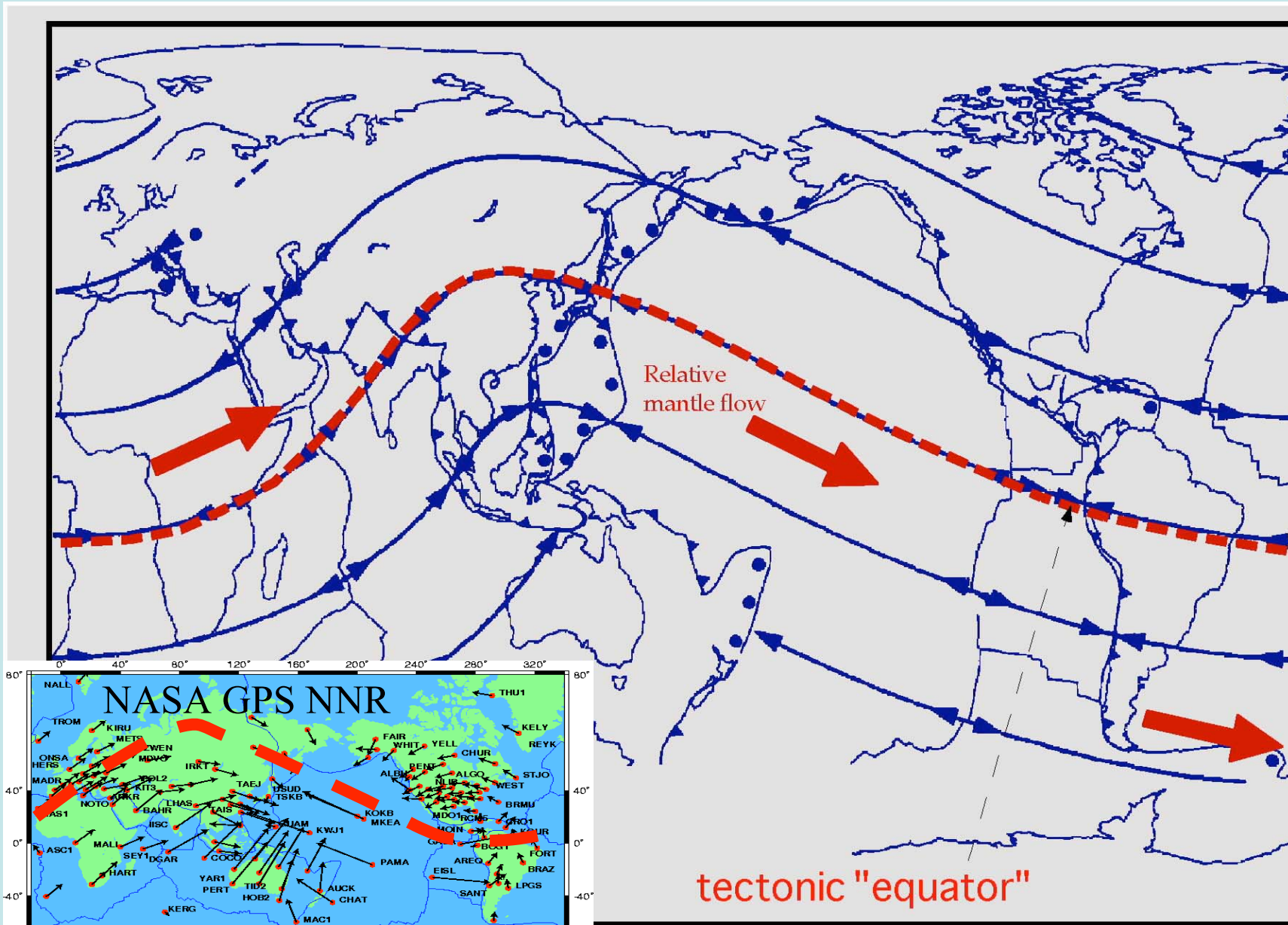


Four subduction zones contributed to deform the area:

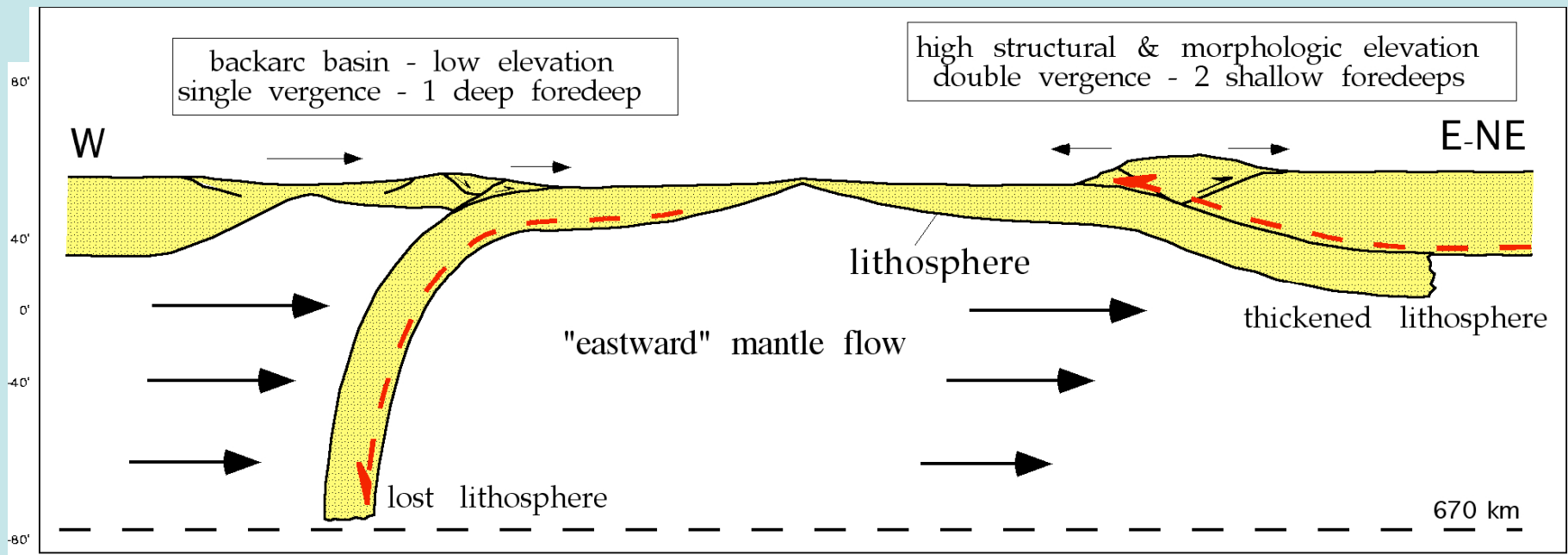
- 1 - ALPS (retrobelt)
- 2 - DINARIDES (forebelt)
- 3 - CARPATHIANS (western backarc)
- 4 - APENNINES (foreland flexure)



- Independent geodynamic processes may coexist in one area
- Contractional belts (e.g., Alps) may form without syn-orogenic collapse



Mainstream of plate motions based on tectonic features



Turcotte & Schubert, 1982

$$w = \frac{2\rho_m\alpha_v(T_m - T_0)}{\rho_m - \rho_w} \left(\frac{kx}{\pi u} \right)^{1/2}$$

W = water depth

ρ_m = asthenosphere density

ρ_w = water density

T_m = mantle temperature

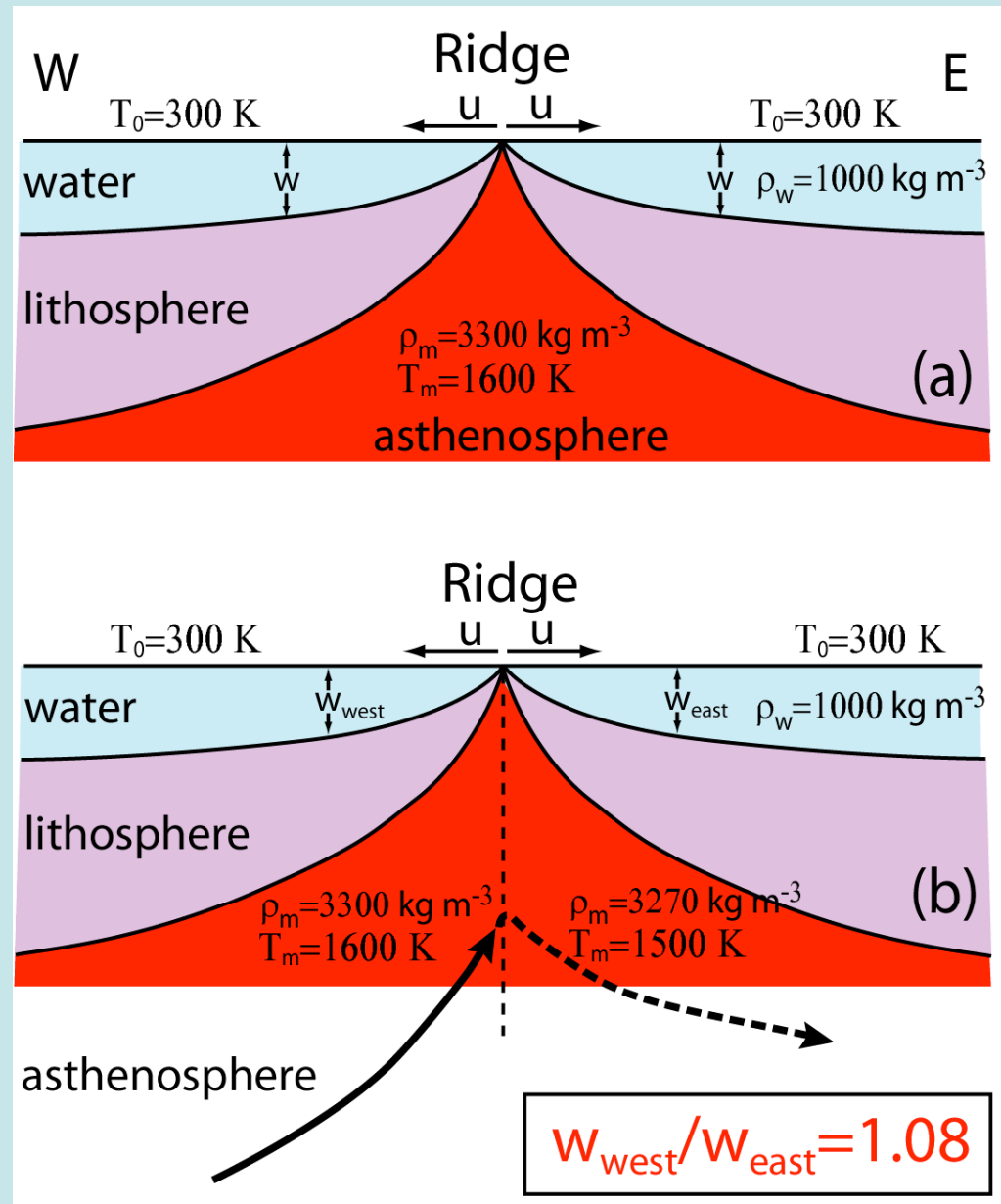
T_0 = surface temperature

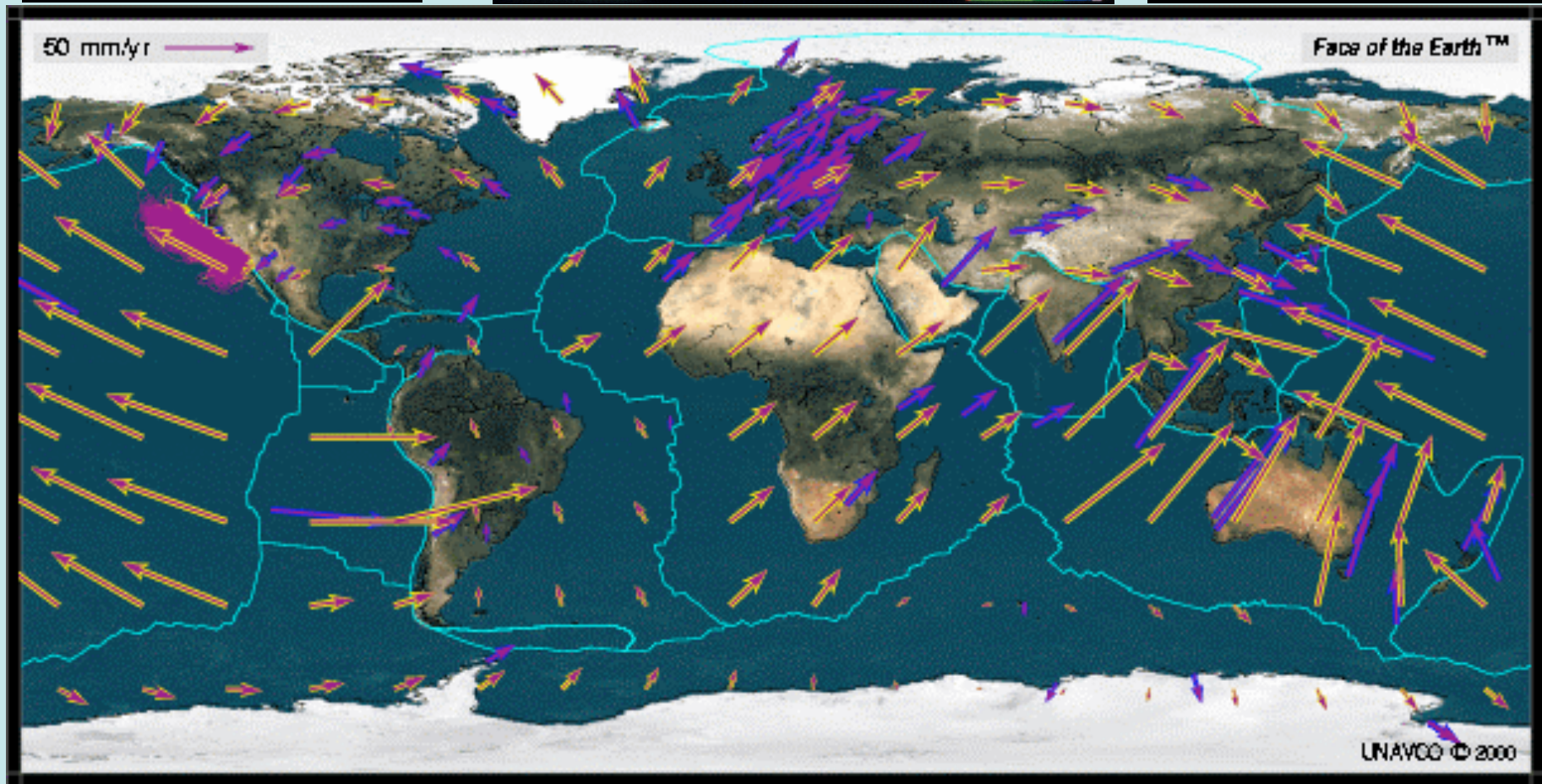
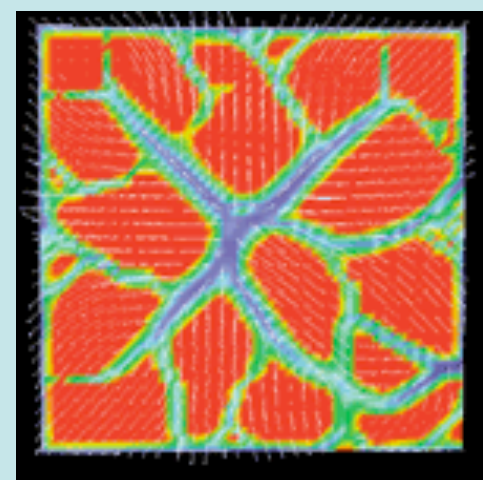
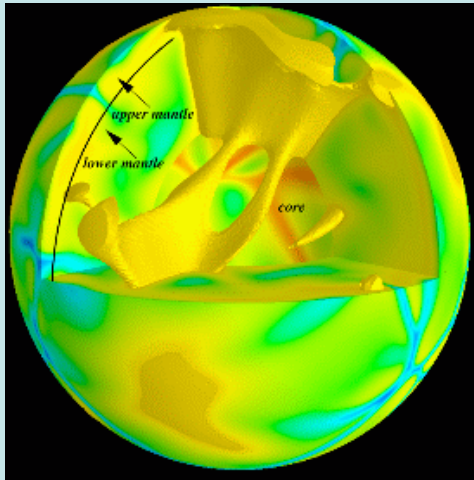
α_v = volumetric coefficient
of thermal expansion

k = mantle thermal diffusivity

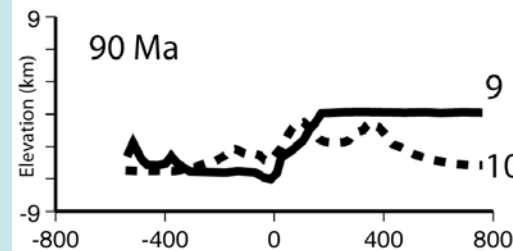
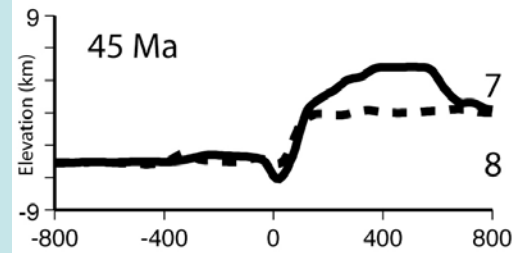
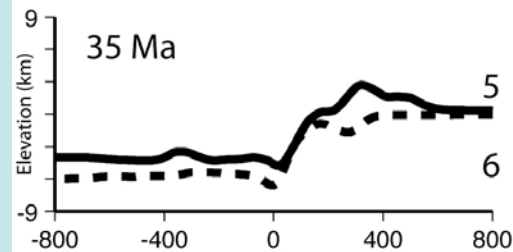
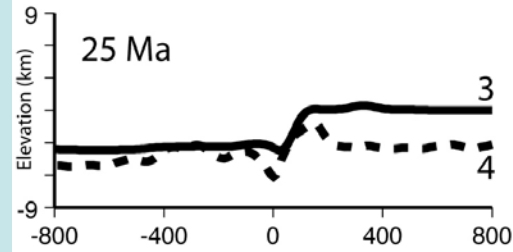
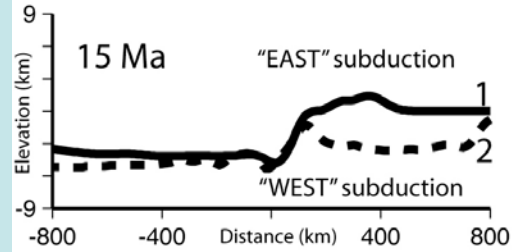
x = distance from the ridge

u = half spreading rate

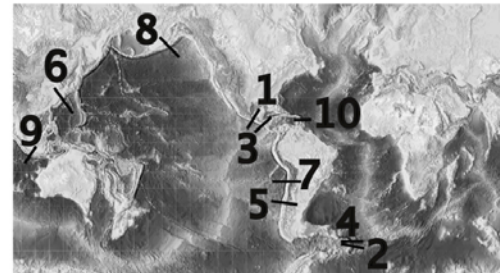
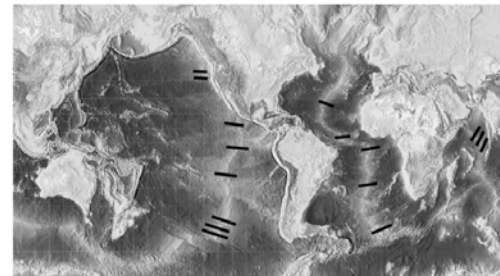
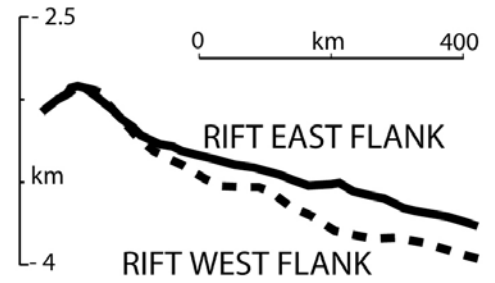


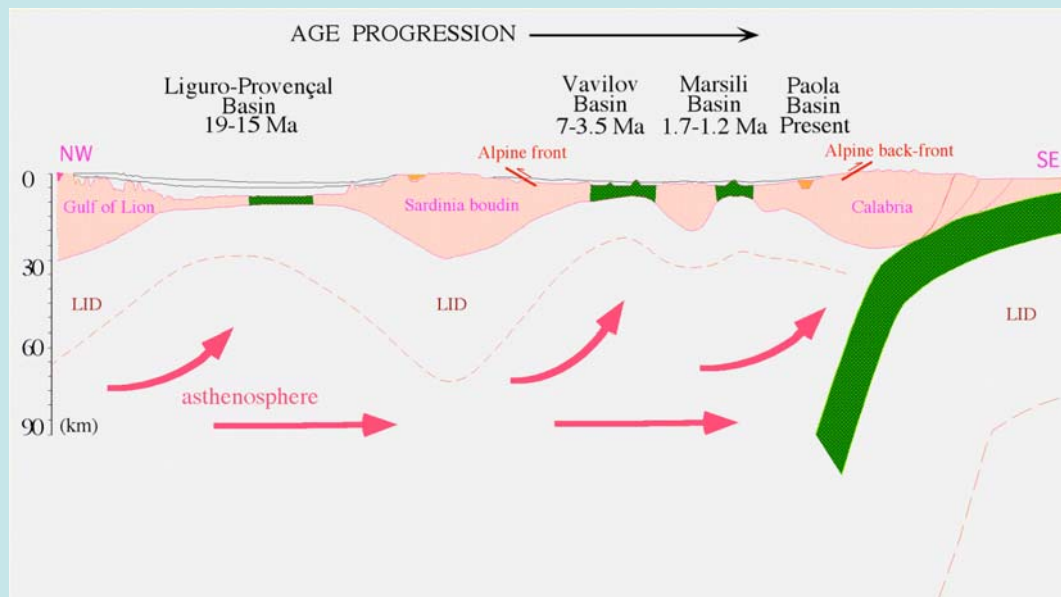
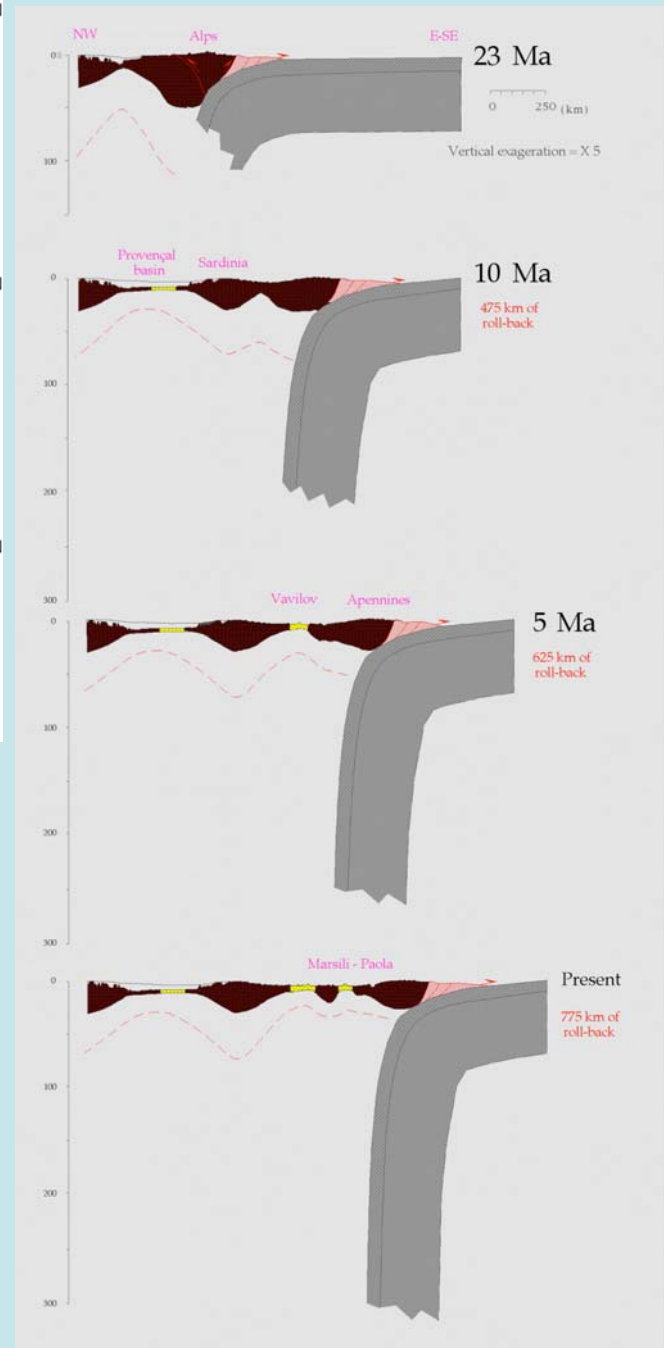
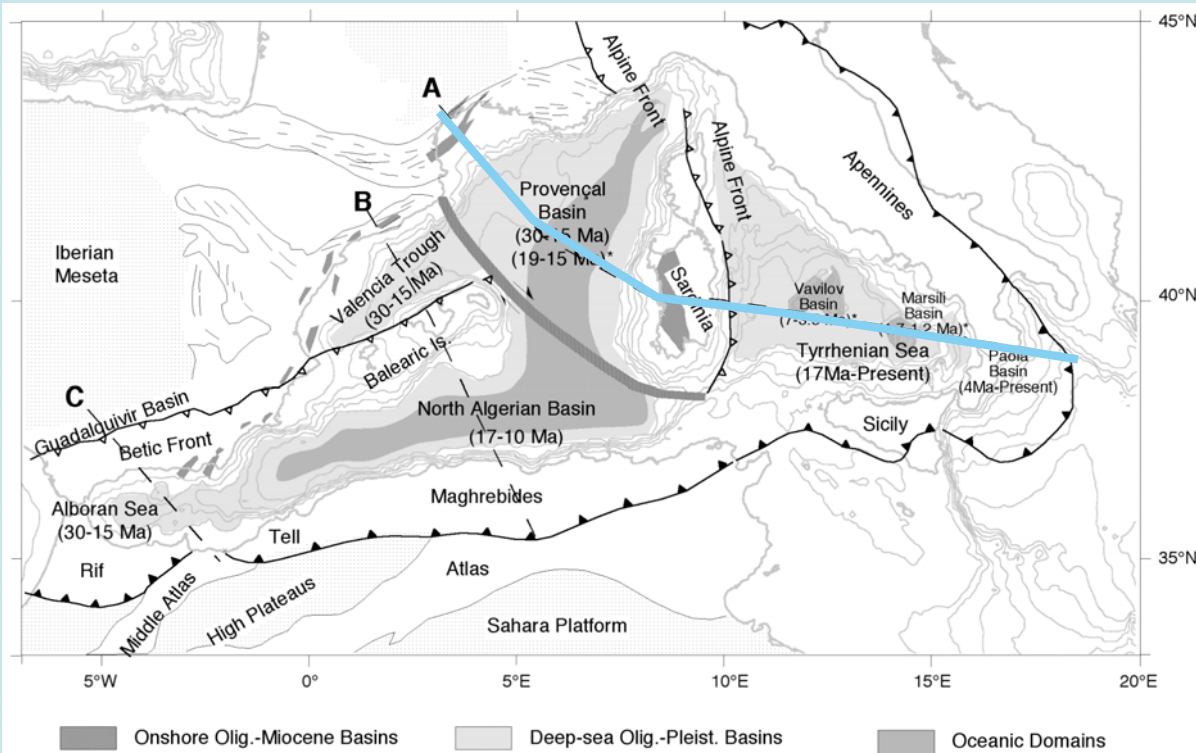


SUBDUCTION ZONES



RIFT ZONES



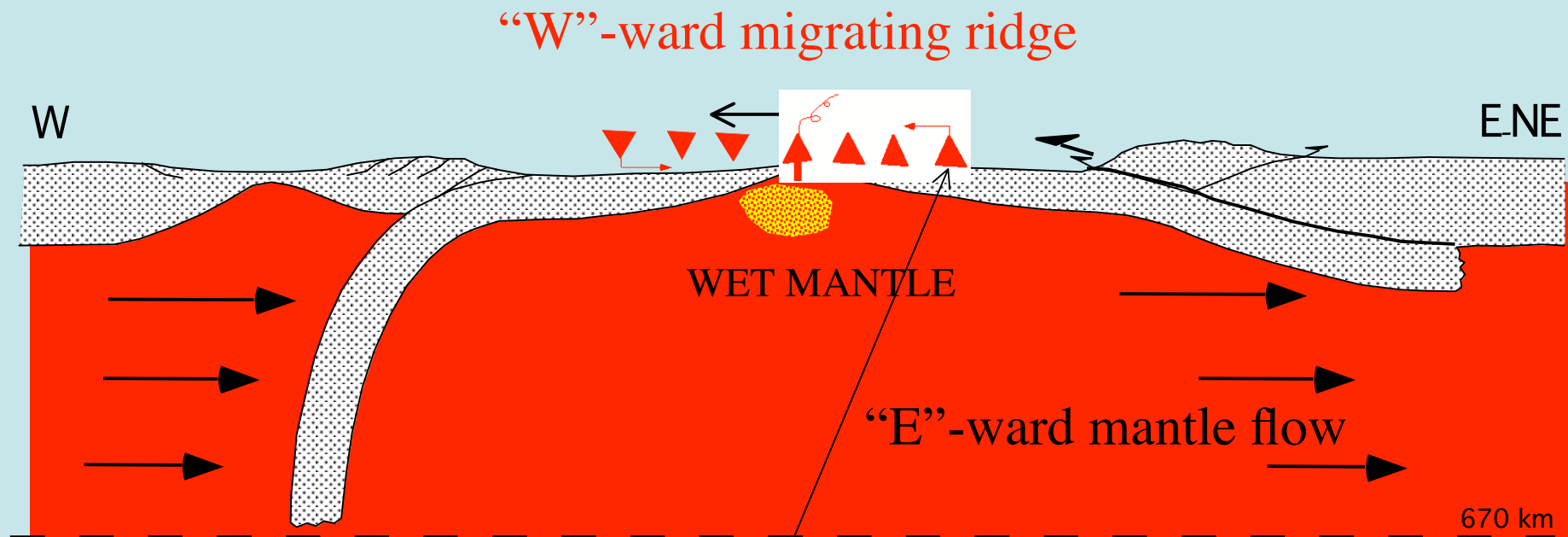


WESTWARD DRIFT OF THE LITHOSPHERE

- Asymmetry of subduction zones
- Asymmetry of rift zones

Evidences of mantle flow also from

- *shear-wave splitting*
- *sheared asthenospheric xenoliths*



W-ward migrating volcanic track

SHALLOW SOURCE OF NOT SO HOT HOTSPOT

Tidal Friction

