



**The Abdus Salam
International Centre for Theoretical Physics**



SMR/1837-5

2007 ICTP Oceanography Advanced School

30 April - 11 May, 2007

Two-Layer Exchange Flows

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Two-Layer Exchange Flows

Lecture for ICTP Advanced School on Oceanography
International Centre for Theoretical Physics
Trieste, Italy April-May 2007

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UNIVERSITY OF SOUTHAMPTON AND
NATURAL ENVIRONMENT RESEARCH COUNCIL



Miller (1983)



Fig. 9.1. The Mediterranean Sea.

Reprinted from *Estuaries and Enclosed Seas*, edited by B.H. Ketchum
 © 1983 Elsevier Scientific Publishing Company, Amsterdam — Printed in The Netherlands

Marsigli (1681)

Fig. II

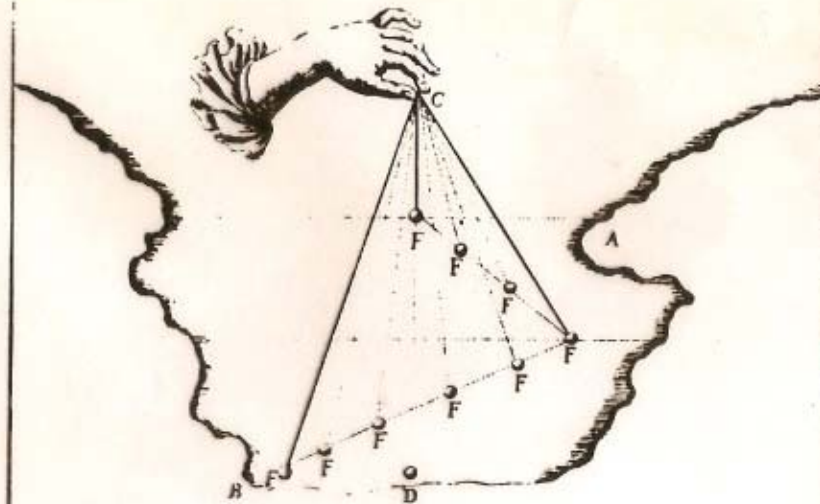
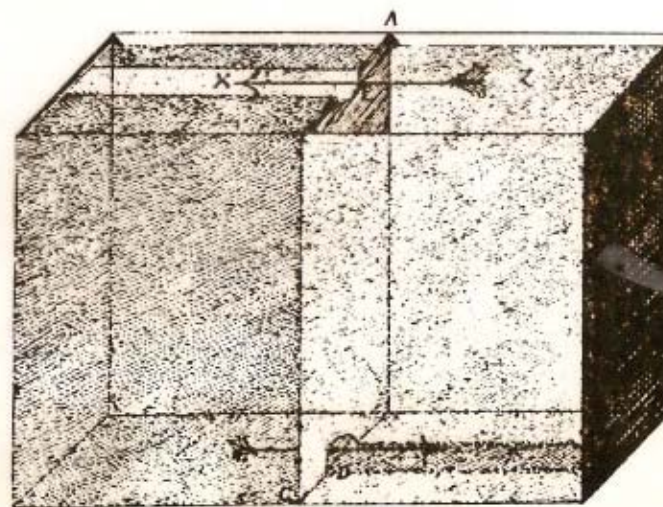
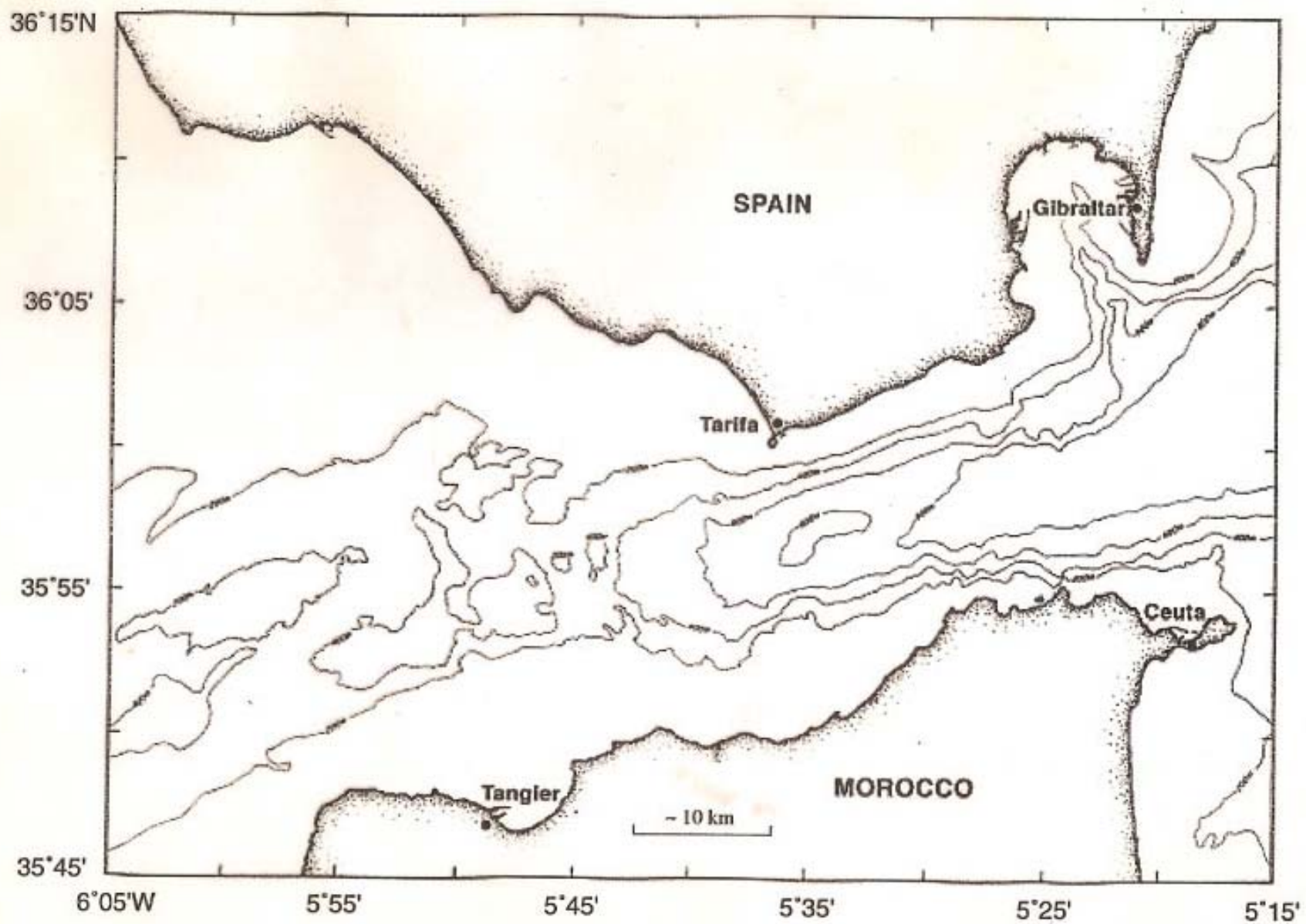
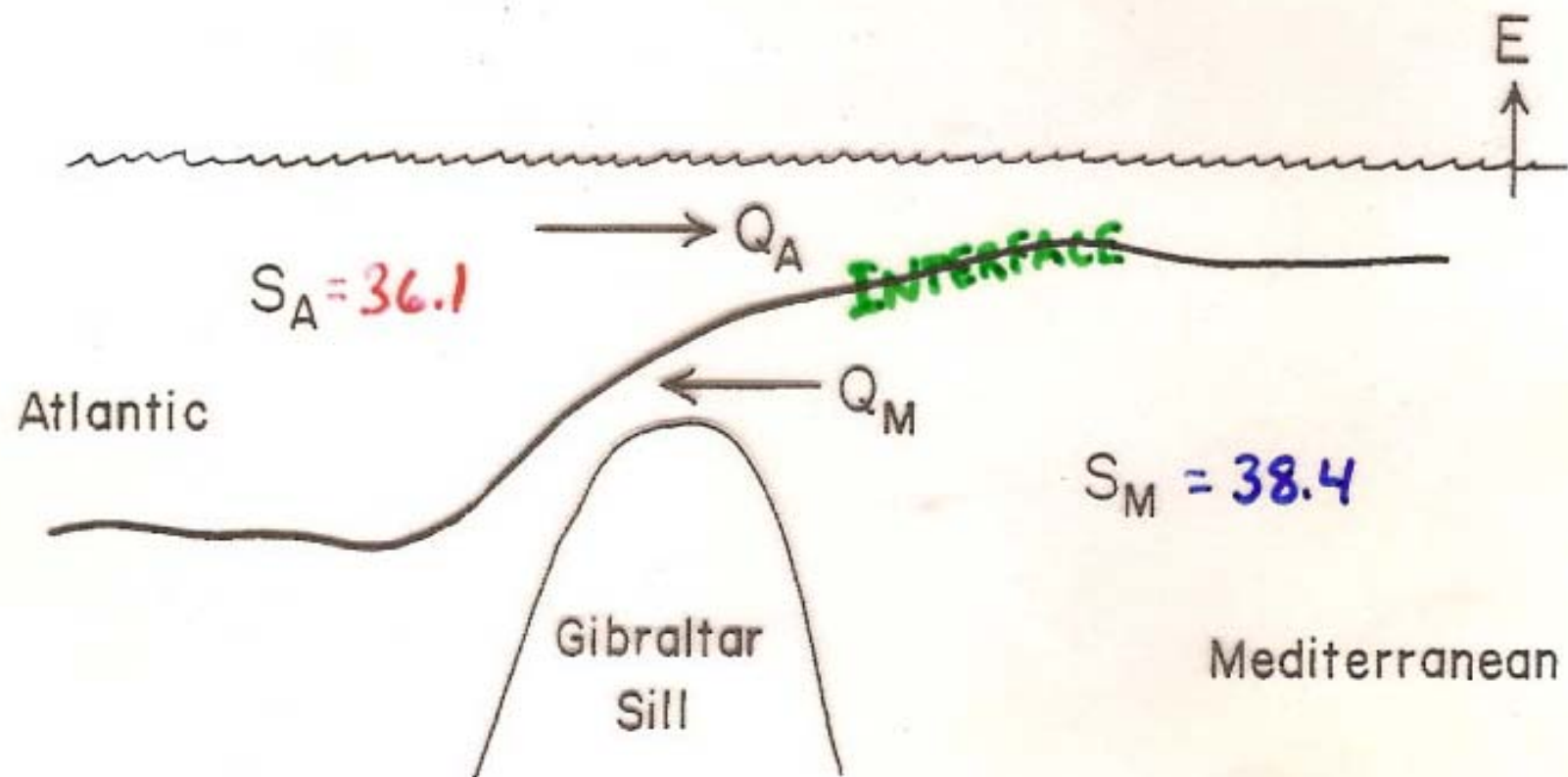
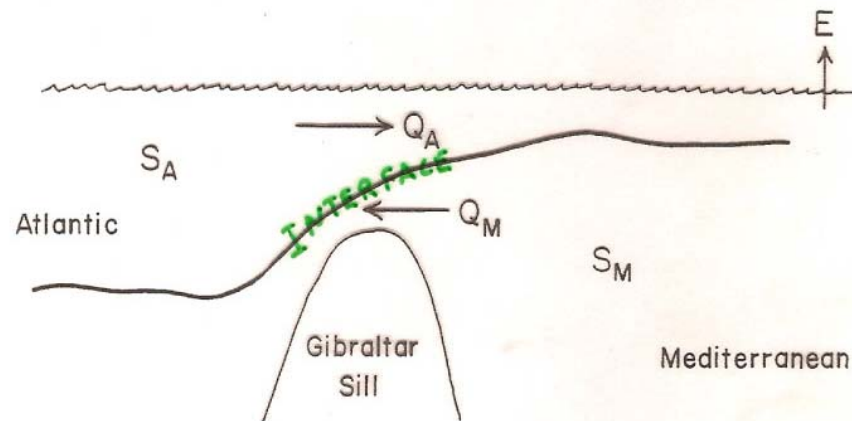


Fig. III









$$Q_A + Q_M = E$$

Mass Conservation

$$Q_A S_A + Q_M S_M = 0$$

Salt Conservation

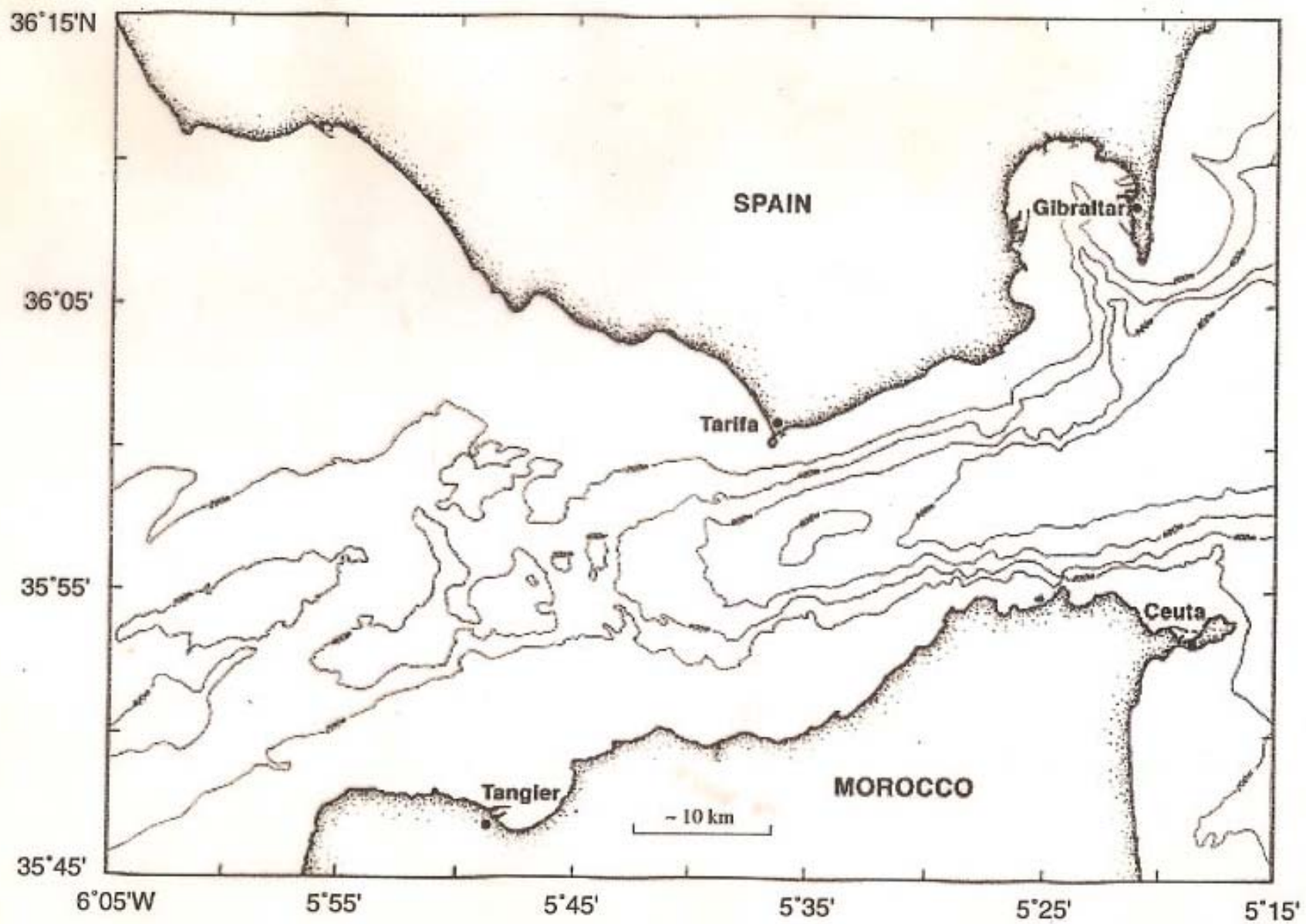
$$Q_A = \frac{E * S_M}{S_M - S_A}$$

Knudsen Relations

$$Q_M = \frac{-E * S_A}{S_M - S_A}$$

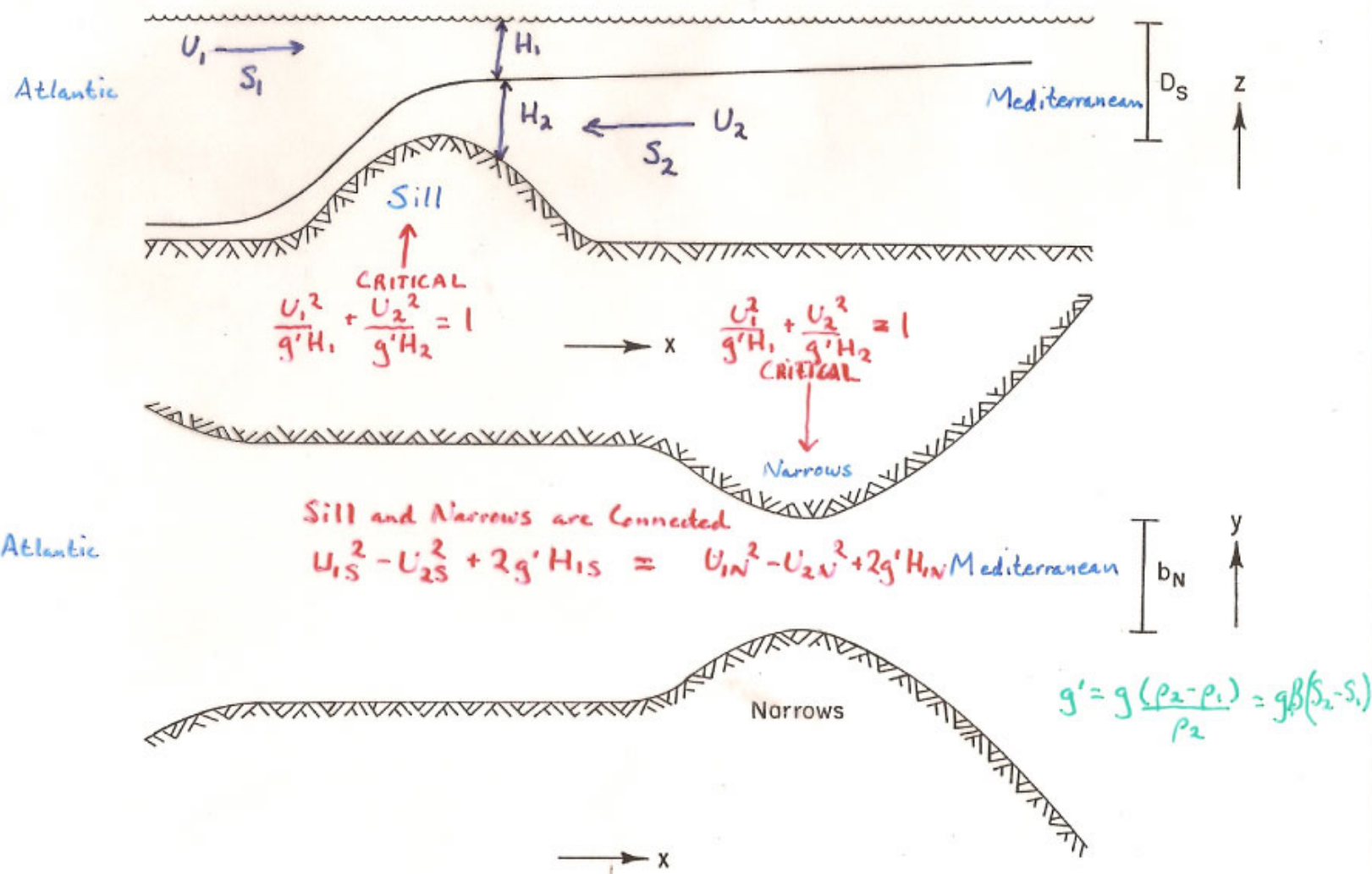
TRADITIONAL METHOD

1. ESTIMATE NET EVAPORATION OVER MEDITERRANEAN
2. MEASURE S_A, S_M IN STRAIT OF GIBRALTAR
3. CALCULATE Q_A, Q_M



HYDRAULIC CONTROL MODEL

NONLINEAR, STEADY, NO FRICTION, NO MIXING, NO ROTATION



Dalziel (1988)

HYDRAULIC CONTROL AND CONSERVATION OF MASS+SALT

$$\underbrace{.277 \sqrt{g' D_s} \frac{D_s b_s}{2}}_{\text{Hydraulic Control}} = \underbrace{Q_1 - Q_2 = \frac{S_M + S_A}{S_M - S_A} E}_{\text{Mass+Salt Conservation}}$$

$$g' = \frac{g \beta (S_M - S_A)}{\rho_0} = \frac{g \beta \Delta S}{\rho_0} \quad S_M = S_A + \Delta S$$

$g, \beta, \rho_0, D_s, b_s, S_A$ are known constants

Solve for ΔS as function of E :

$$\frac{\Delta S^3}{\left(1 + \frac{\Delta S}{2 S_A}\right)^2} = \frac{\rho_0}{g \beta D_s} \left(\frac{2 S_A E}{.277 \frac{D_s b_s}{2}} \right)^2$$

Mass + Salt Conservation suggest

$$|Q_A| + |Q_M| \propto \frac{E}{S_M - S_A}$$

Hydraulic Control Theory suggest

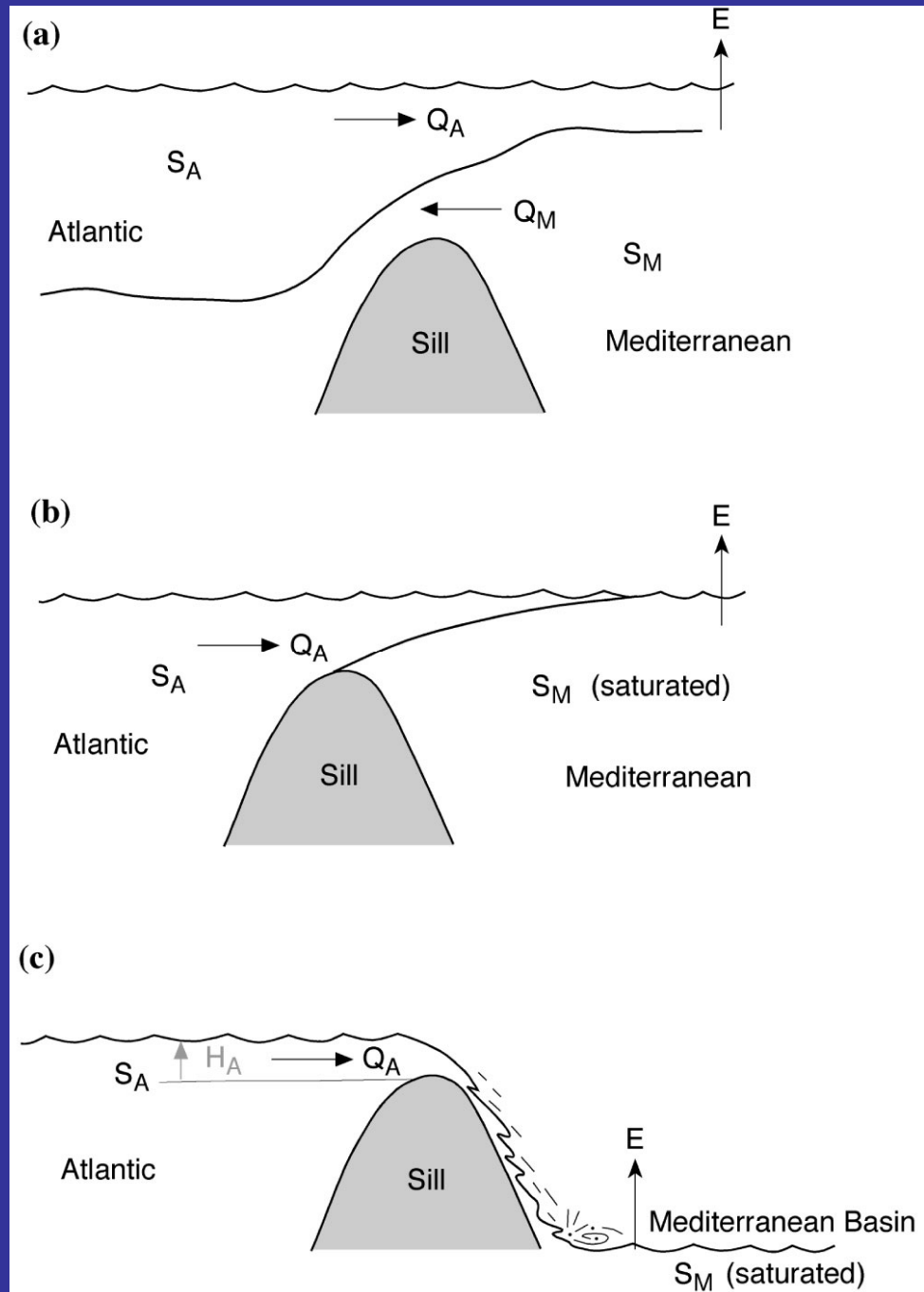
$$|Q_A| + |Q_M| \propto \sqrt{S_M - S_A}$$

For both to be satisfied

$$S_M - S_A \propto E^{2/3}$$

Table 1. Prediction of the salinity difference, inflow and outflow transports for the steady, two-layer maximal exchange through the Strait of Gibraltar as a function of the net evaporation over the Mediterranean Basin. The net evaporation over the Mediterranean Basin, E , equals the average local net evaporation, e , times the surface area of the Mediterranean, $2.52 \times 10^{12} \text{ m}^2$

Net evaporation $e \text{ (cm y}^{-1}\text{)}$	Salinity difference, $\Delta S = S_2 - S_1 \text{ (‰)}$	Inflow transport, $Q_1 \text{ (10}^6 \text{ m}^3 \text{ s}^{-1}\text{)}$	Outflow transport, $-Q_2 \text{ (10}^6 \text{ m}^3 \text{ s}^{-1}\text{)}$
50	1.75	0.87	0.83
60	1.98	0.92	0.88
70	2.20	0.98	0.92
80	2.40	1.03	0.96
90	2.61	1.07	1.00
100	2.80	1.11	1.03



THE FILLING OF THE MEDITERRANEAN SEA

BY GUY BILLOUT

