



SMR/1849-4

### Conference and School on Predictability of Natural Disasters for our Planet in Danger. A System View; Theory, Models, Data Analysis

25 June - 6 July, 2007

Possibilities opened by new data bases - GPS

> G.F. PANZA Department of Earth Sciences/ICTP Trieste



The evaluation of seismic hazard is based on the traditional Probabilistic Seismic Hazard Analysis, i.e. on the probabilistic analysis of earthquake catalogues and of ground motion, from macroseismic observations and instrumental recordings. This leads to severe bias in the estimation of seismic hazard, with artificially inflated errors, because the mathematical model of PSHA, as it is in use today, is inaccurate and leads to systematic errors in the calculation process.

Recently this approach showed its limitation in providing a reliable seismic hazard assessment, possibly due to the insufficient information about historical seismicity, which can introduce relevant errors in the purely statistical approach mainly based on the seismic history. The probabilistic analysis supplies indications that can be useful but are not sufficiently reliable to characterize seismic hazard. Recent examples Kobe (17.1.1995), Bhuj (26.1.2001), Boumerdes (21.5.2003) and Bam (26.12.2003) events.

<b>? GSHAP ?</b> Kobe (17.1.1995), Gujarat (26.1.2001), Boumerdes (21.5.2003) and Bam (26.12.2003) earthquakes PGA(						
with a in 50	probability of exceedence of 109 ) years (return period 475 years)	6				
Kobe	0.40-0.48	0.7-0.8				
<ul> <li>Gujarat</li> </ul>	0.16-0.24	0.5-0.6				
Boumerdes	0.08-0.16	0.3-0.4				
• Bam	0.16-0.24	0.7-0.8				

To overcome the mentioned limitations and, above all, to improve the preseismic information which may lead to an effective mitigation of seismic risk, we are following an innovative that Earth approach, combines Observation (EO) data and new advanced approaches in seismological and geophysical data analysis.



The system we are developing is based on the neodeterministic approach for the estimation of seismic ground motion, integrated with the space and time dependent information provided by EO data analysis through geophysical forward modeling.

The need of integration of different geophysical observables is obvious when the process of earthquake preparation and occurrence is analysed: the lithosphere - a hierarchical system of interacting blocks - accumulates stress, according to strain and strain rates fields due to tectonics, which is partly released during the earthquake occurrence.



## Seismological data analysis

- Data on seismicity (earthquake catalogues), geomorphology and geodynamics and Earth structure (velocity, gravity data);
- Worldwide tested pattern recognition algorithms for middle-range intermediateterm earthquake prediction and for identification of damaging earthquake prone areas;
- Robust and tested codes for the earth structure retrieval and numerical modelling of lithosphere block dynamics.

### Seismological data analysis

- OUTPUT (1) Regional alerted areas by the near real time monitoring of seismicity (TIPs for the occurrence of earthquakes with  $M \ge M_0$ );
- Maps of the morphostructural zonation and selection of seismogenic nodes prone to earthquakes with M≥6.0 & M≥6.5 within the regional alerted regions;

### Seismological data analysis • OUTPUT (2)

- Restrained local alerted areas for GPS and SAR investigations;
- Multiscale velocity models of the Earth Structure for geophysical forward modelling;

Preferred models for the dynamics of the lithosphere at a regional scale.





# Intermediate-term middle-range earthquake prediction experiment

CN algorithm (Keilis-Borok et al., 1990; Peresan et al., 2005) M8S algorithm (Kossobokov et al, 2002)

### **Main features:**

- Fully formalized algorithms and computer codes available for independent testing;
- Use of published & routine catalogues of earthquakes;
- Worldwide tests ongoing for more than 10 years permitted to assess the significance of the issued predictions (Kossobokov et al., 1999; Rotwain and Novikova, 1999)

### Intermediate-term middle-range earthquake prediction experiment in Italy CN and M8S algorithms are based on a set of empirical functions of time to allow for a quantitative analysis of the premonitory patterns which can be detected in the seismic flow: Variations in the seismic activity Seismic quiescence Space-time clustering of events They allow to identify the TIPs (Times of Increased Probability) for the occurrence of a strong earthquake within a delimited region

### Intermediate-term middle-range earthquake prediction experiment in Italy

- Stability tests with respect to several free parameters of the algorithms (e.g. Costa et al., 1995; Peresan et al., GJI, 2000; Peresan et al., PEPI, 130, 2002);
- CN predictions are regularly updated every two months since January 1998;
- M8S predictions are regularly updated every six months since January 2002;

Real time prediction experiment started in July 2003



Experiment	Space-time volume of alarm (%)	n/N	Confidence level (%)
Retrospective* (1954 – 1963)	41	3/3	93
Retrospective (1964 - 1997)	27	5/5	>99
Forward (1998 – 2007)	36	4/5	94
All together (1954 – 2007)	31	12/13	>99
ntral and Southern reg	ons only		

Experiment	M6.5+		M6.0+		M5.5+	
	Space- time	n <i>i</i> N	Space- time	n/N	Space- time	n/N
Retrospective (1972-2001)	36	2/2	40	1/2	39	9/14
Forward (2002-2007)	49	0,0	43	0,0	25	5/9
All together (1972-2007)	37	2/2	40	1/2	38	14/2









• The Morphostructural Zonation method, MSZ (*Alekseevskaya et al.,* 1977), allows to identify, independently from seismicity information, the sites where strong earthquakes are likely to occur.







 Maps of areas alerted by CN and M8s will be compared with EO information, taking into account modelling of the reology provided by Geophysical Modelling;





• Stress maps at the depth of the active faults will be obtained through integration of EO geodetic information into Geophysical Forward Modelling.



















#### NEODETERMINISTIC

New approach based on synthetic signals

attenuation







































130° 135° 140°

Number of earthquakes occurred in 2-month intervals, within (a) and outside (b) the snowy region. Red and white histograms show M≥7.0 (left) and 7.0>M≥6.0 (right) events, respectively. In (c) blue squares show maximum snow depths in a winter at AMeDAS stations (only points with snows deeper than 20.0 cm are shown). Epicenters of M≥7.0 earthquakes are shown in (c) as circles (snowy region) and triangles (outside). Red curve in (a) is the best-fit probability density function of the earthquake occurrence based on the twocomponent (stationary and annual) model (Heki, EPSL, 2003).











Major events in Himalaya				
~1100 (>8.5)	1255 (7.5-8.0)?			
~1413 (>8.5)	1505 (>8.5)			
1555 (8.0-8.5)	1681 (7.5-8.0)?			
1724 (7.5-8.0)	1803 (7.5-8.0)			
1833 (7.5-8.0)	1897 (7.5-8.0)			
1905 (7.5-8.0)	1934 (8.0-8.5)			
1947 (7.5-8.0)	1950 (>8.5)			
2005 (7.5-8.0)				
	(Upreti, 2007)			





